



29 **1 Introduction**

30 Despite accounting for only a small fraction of the global freshwater inventory, rivers
31 constitute one of the most dynamic components of the hydrological cycle (Oki and Kanae, 2006;
32 Palmer and Ruhi, 2018). They not only provide the most accessible freshwater resource for human
33 societies but also serve as critical biogeochemical conduits linking terrestrial and marine
34 ecosystems through the transport of sediments, nutrients, and thermal energy (Syvitski et al., 2005;
35 Allen and Pavelsky, 2018; Resplandy et al., 2018). Consequently, the accurate and continuous
36 quantification of river discharge is fundamental for sustainable water resource management,
37 reliable hydroclimatic modeling, and assessments of environmental change.

38 Traditionally, in situ gauging networks have served as the primary benchmark for river
39 discharge monitoring, providing high-fidelity and temporally continuous observations (Alsdorf et
40 al., 2007; Hannah et al., 2011; Andrews and Grantham, 2024). However, the global distribution of
41 these networks is highly uneven. While North America and Europe maintain relatively dense
42 monitoring systems, extensive regions across Asia, Africa, and South America remain poorly
43 gauged or entirely data-sparse (Do et al., 2018; Krabbenhoft et al., 2022; Collins et al., 2024; Oh
44 and Bartos, 2025; Sun et al., 2026). Moreover, even where gauging infrastructure exists, the access
45 to discharge records is often restricted by geopolitical barriers and limited data-sharing policies
46 (Fekete et al., 2015; Famiglietti et al., 2015; Ruhi et al., 2018; Collins et al., 2024). Together, these
47 limitations substantially hinder the global characterization of fluvial dynamics and have motivated
48 the rapid development of satellite-based Earth observation and hydrological modeling approaches
49 to address observational gaps (Gleason et al., 2014; Lin et al., 2019; Ghiggi et al., 2019; Hao et al.,
50 2024; Scherer et al., 2024).

51 Hydrological models can provide spatially and temporally continuous simulations of river
52 discharge, yet their predictive accuracy remains constrained by uncertainties in meteorological
53 forcing data and simplified representations of hydrological processes (Beven, 2007; Zheng et al.,
54 2019; Tang et al., 2023). In contrast, satellite remote sensing has emerged as a promising
55 alternative, owing to its unprecedented global coverage and continually improving spatiotemporal
56 resolutions (Lettenmaier et al., 2015; Sichangi et al., 2017; Hou et al., 2020). Although satellites
57 cannot directly measure river discharge, they can routinely retrieve key hydraulic variables such
58 as water surface elevation, river width, and surface slope (Smith et al., 1997; Bonnema et al., 2016;
59 Durand et al., 2016). When integrated with historical in situ measurements, these remotely sensed



60 variables can be converted into continuous discharge estimates using empirical hydraulic geometry
61 relationships and rating-based approaches (Pavelsky, 2014; Paris et al., 2016). Conceptually, this
62 framework parallels conventional stage–discharge rating curves employed at gauging stations,
63 with satellite-derived river width or water surface elevation replacing ground-based stage
64 measurements.

65 In recent years, satellite earth observations have facilitated substantial progress in extending
66 global river discharge records. Notably, the Satellite Altimetry-based Extension of global-scale in
67 situ river discharge Measurements (SAEM) integrated multi-mission satellite altimetry with
68 existing stream gauging networks to reconstruct discharge at 8,730 global stations, representing
69 approximately 88% of the globally gauged discharge volume (Saemian et al., 2024). Meanwhile,
70 optical remote sensing has demonstrated strong capability for width-based discharge estimation.
71 Riggs et al. (2023) utilized Landsat and Sentinel-2 imagery to extract dynamic river widths and
72 extended discharge records for 2,168 gauges globally. Similarly, Lin et al. (2023) conducted a
73 large-sample evaluation of width-based discharge algorithms across 3,078 gauges, demonstrating
74 their scalability and sensitivity to hydraulic and geomorphological conditions. More recently,
75 multi-sensor integration has emerged as a promising strategy to maximize observational coverage
76 and reliability. Elmi et al. (2024) combined satellite-derived river widths with altimetry-based
77 water levels and applied a stochastic non-parametric framework to reconstruct monthly discharge
78 time series for 3,377 discontinued Global Runoff Data Centre (GRDC) stations, resulting in the
79 Remote Sensing-based Extension for the GRDC (RSEG) dataset with rigorous uncertainty
80 quantification.

81 Despite these substantial global advances, a critical geographic gap persists in China. China
82 hosts one of the world’s largest and most diverse river systems and is undergoing profound
83 hydrological alterations. River discharge regimes across the country are increasingly shaped by
84 the combined influences of climate change—including accelerated glacier retreat over the Tibetan
85 Plateau—and intensified anthropogenic regulation such as land use change and reservoir operation
86 (Yao et al., 2022; Yang et al., 2022; Wang et al., 2025a; Wang et al., 2025b). Nevertheless, due to
87 restrictive data-sharing policies, high-frequency in situ discharge observations from China remain
88 largely inaccessible. This scarcity is particularly evident in the GRDC, the most comprehensive
89 global archive of river discharge observations, which currently contains records for only 31 gauges
90 within China (Färber et al., 2025). In addition, China is severely underrepresented in existing

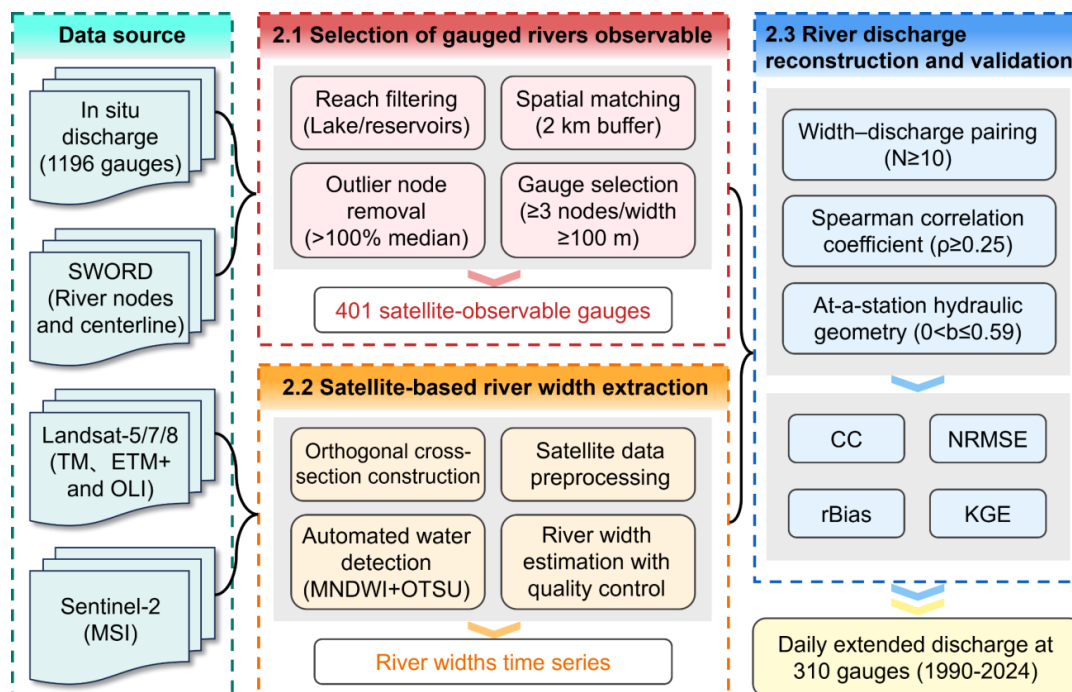


91 global remote sensing discharge products. For example, the SAEM dataset encompasses only 62
92 gauges across China, accounting for less than 1% of its total reconstructed gauges (Saemian et al.,
93 2024). This stark disparity underscores a major deficiency in the global characterization of river
94 discharge dynamics across one of the world's most hydrologically important regions.

95 To address this gap, this study leverages river width observations derived from Landsat and
96 Sentinel optical imagery to establish station-specific width–discharge relationships and reconstruct
97 daily river discharge records for 310 gauges across China. To our knowledge, this represents the
98 most comprehensive satellite-derived discharge reconstruction for China to date. Compared with
99 existing global datasets, the number of gauge-constrained stations within China has increased by
100 at least fivefold. Across all validation gauges, the reconstructed discharge achieves a median
101 Kling–Gupta Efficiency (KGE) of 0.66, outperforming currently available global products. The
102 resulting dataset, termed the China Daily River Discharge Records (CDR²), substantially expands
103 long-term discharge observations in a hydrologically critical yet data-sparse region. Beyond
104 improving regional hydrological monitoring, CDR² provides an essential benchmark for assessing
105 riverine responses to climate change and anthropogenic disturbances and contributes valuable
106 observations for future global hydrological modeling and satellite calibration efforts.

107 **2 Data and Methods**

108 This section describes the datasets, remote sensing products, and methodological framework
109 used to reconstruct and extend river discharge observations across China. The workflow integrates
110 in situ discharge measurements, satellite-derived river width observations from Landsat and
111 Sentinel-2 imagery, and station-specific hydraulic geometry relationships to generate continuous
112 daily discharge records. Fig. 1 provides an overview of the technical framework, including data
113 acquisition, river width extraction, construction of width–discharge relationships, discharge
114 reconstruction, and validation procedures. Detailed descriptions of the datasets, preprocessing
115 steps, and reconstruction methods are presented in the following subsections.



116

117 **Figure 1.** Technical framework of the study.

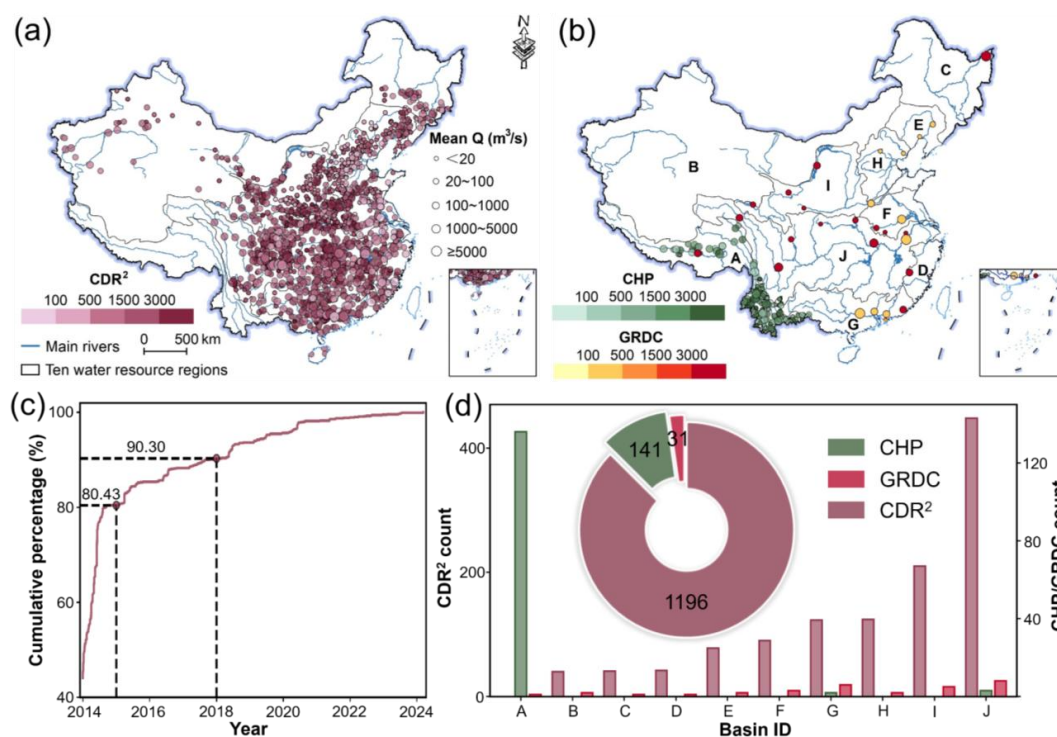
118 **2.1 Selection of gauged rivers observable**

119 Daily in situ river discharge observations were obtained from the National Hydrological and
 120 Rainfall Information System of the Ministry of Water Resources of China, which integrates
 121 national and provincial monitoring networks spanning major rivers, lakes, and reservoirs and
 122 provides real-time and historical records of water level, discharge, and storage. We compiled daily
 123 discharge records from 1,196 gauging stations across 2014–2024 to construct the CDR² dataset
 124 (Fig. 2a). To minimize redundancy, when multiple gauges were located within 100 m of one
 125 another, only the station with the longest observational record was retained. The resulting gauge
 126 network spans nearly all major river basins in China, except the Southwest River Basin. Mean
 127 discharge among the selected gauges ranges from 0.5 to 30,000 m³/s, with each station containing
 128 an average of 1,945 observations. In terms of temporal coverage, more than 80% of the gauges
 129 contain observations beginning in or before 2015, while over 90% include records extending back
 130 to 2018 or earlier (Fig. 2c). Collectively, these characteristics demonstrate that CDR² provides



131 both extensive spatial coverage and high-frequency discharge observations across China, offering
 132 a valuable foundation for large-scale hydrological reconstruction and analysis.

133 Compared with existing publicly available global discharge datasets, CDR² substantially
 134 improves both the spatial coverage and observational availability of river discharge records across
 135 China. The GRDC, currently the most comprehensive global archive of in situ discharge
 136 observations, contains records from only 31 gauges in China, many of which predate 2004 and
 137 have relatively short observational periods, including some with only a single year of data.
 138 Similarly, the Chinese Hydrology Project (CHP) dataset provides discharge records from 141
 139 gauges, primarily spanning 1953–1987 (Henck et al., 2010; Schmidt et al., 2011). In contrast,
 140 CDR² incorporates approximately 8.5 times as many gauges as CHP and 38.6 times as many as
 141 GRDC (Fig. 2d), while also offering substantially more recent observations at a daily temporal
 142 resolution.

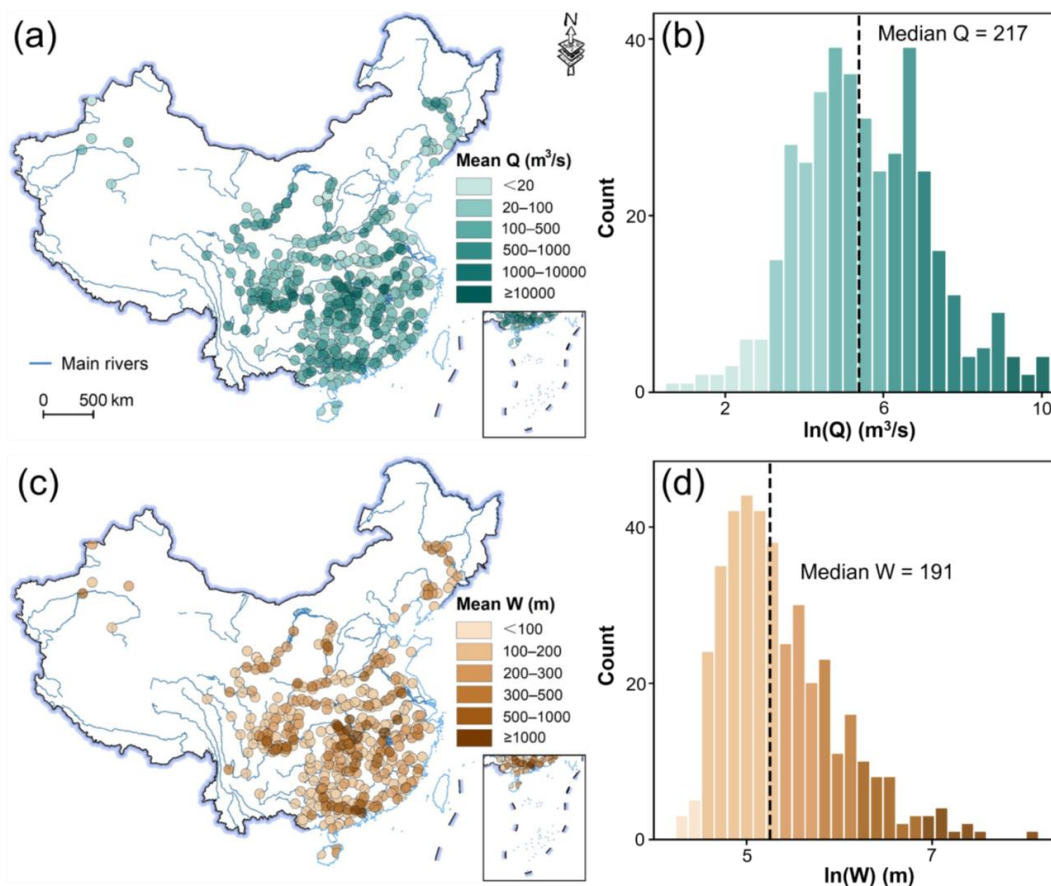


143
 144 **Figure 2.** Spatial distribution and temporal coverage of in situ river discharge observations across
 145 China (2014–2024). (a) Spatial distribution of the 1,196 gauges included in the China Daily River
 146 Discharge Records (CDR²) dataset, with symbol size representing mean discharge (Q, m³/s). (b)
 147 Spatial distribution of river gauges available from the Chinese Hydrology Project (CHP) and the



148 Global Runoff Data Centre (GRDC) within China. The labeled regions (A–J) denote the ten major
149 basins: Southwest Rivers (A), Northwest Rivers (B), Songhua River (C), Southeast Rivers (D),
150 Liao River (E), Hai River (F), Pearl River (G), Huai River (H), Yellow River (I), and Yangtze
151 River (J). **(c)** Cumulative probability distribution of gauge observation start years in the CDR²,
152 indicating that over 80% of gauges contain records beginning in or before 2015, while more than
153 90% begin in or before 2018. **(d)** Comparison of the number of gauges included in the CDR², CHP,
154 and GRDC datasets across China.

155 To identify gauges suitable for satellite-based discharge reconstruction, we employed the
156 Surface Water and Ocean Topography River Database (SWORD) (Altenau et al., 2021) as the
157 reference river network. SWORD provides a globally consistent river centerline dataset with
158 connected topology, in which river systems are segmented into predefined reaches and nodes. To
159 minimize the influence of non-channel water bodies, river reaches linked to lakes and reservoirs
160 were excluded using the SWORD lakeflag attribute. For each gauge, all river nodes located within
161 a 2 km buffer radius were extracted to establish the spatial correspondence between in situ
162 discharge observations and satellite-observed river width. To improve robustness, outlier nodes
163 were removed using a median-based relative deviation criterion, whereby nodes with width
164 deviations exceeding 100% of the median width were considered anomalous and excluded. A
165 gauge was retained only if it satisfied two conditions: (1) the presence of at least three valid
166 matched nodes, and (2) at least one node with a mean river width of ≥ 100 m, as defined in SWORD.
167 These criteria ensured that the river channel could be reliably detected by optical satellite imagery
168 and that the extracted width measurements were sufficiently representative for discharge
169 reconstruction. Following the screening procedure, 401 gauges were identified as suitable for
170 satellite-based analysis (Fig. 3). The mean discharge of these gauges ranges from 1.66 to 27,797.46
171 m³/s, with a median of 217 m³/s and a mean of 1,090 m³/s. The corresponding river widths range
172 from 69 to 3,504 m, with a median of 191 m and a mean of 277 m.



173
174 **Figure 3.** Spatial distribution and hydrological characteristics of gauges selected for satellite-based
175 discharge reconstruction. (a) Spatial distribution of the 401 selected gauges, with symbol color
176 representing mean discharge (Q , m^3/s). (b) Frequency distribution of $\ln(Q)$ for the selected gauges,
177 showing a median discharge of $217 \text{ m}^3/\text{s}$. (c) Spatial distribution of the corresponding mean river
178 width (W , m) for the selected gauges. (d) Frequency distribution of $\ln(W)$, with a median river
179 width of 191 m .

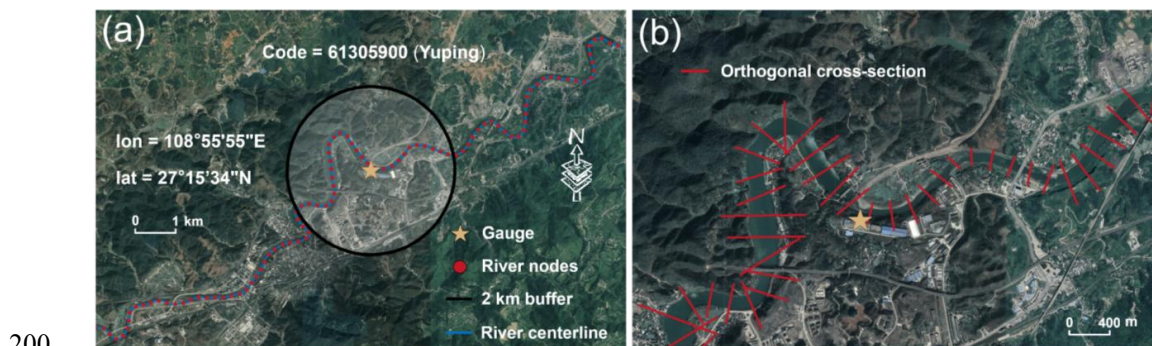
180 2.2 Satellite-based river width extraction

181 2.2.1 Construction of orthogonal river cross-sections

182 To retrieve long-term river width time series, we implemented an automated fixed-geometry
183 approach based on predefined orthogonal cross-sections (Fig. 4). Unlike approaches that
184 dynamically determine cross-section orientation for each individual image, the orthogonal



185 directions in this study were derived once from the river centerline and consistently applied
186 throughout the entire satellite image archive, thereby substantially improving computational
187 efficiency (Pavelsky and Smith, 2008; Yang et al., 2019; Feng et al., 2019). For each matched river
188 node identified in Section 2.1, its location was projected onto the river centerline to determine its
189 along-channel position. Local flow direction was then estimated using two control points sampled
190 symmetrically upstream and downstream along the centerline. To ensure directional stability
191 across rivers of varying sizes, the sampling distance was constrained to 1–10 m. The resulting
192 tangent vector was normalized and rotated by 90° to derive the unit normal vector perpendicular
193 to the local flow direction. Orthogonal cross-sections were centered at the projected node location
194 and extended $1.5\bar{W}$ on both sides of the channel centerline, where \bar{W} represents the prior mean
195 river width derived from SWORD, yielding a total transect length of $3\bar{W}$. This proportional design
196 ensures full coverage of the active river channel while minimizing contamination from adjacent
197 floodplains, wetlands, or nearby water bodies. In cases where local centerline curvature or
198 geometric discontinuities prevented reliable tangent estimation, the orthogonal cross-sections were
199 defined using the perpendicular direction of the nearest valid centerline segment.



200
201 **Figure 4.** Construction of orthogonal cross-sections for satellite-based river width extraction. (a)
202 Example at the Yuping gauge illustrating the gauge location, river centerline, and associated river
203 nodes. (b) Orthogonal cross-sections generated at each river node along the river centerline for
204 extracting river width from satellite imagery. Source of the underlying image: © Google Earth.

205 2.2.2 Satellite data preprocessing and surface water detection

206 River width was derived from optical satellite imagery acquired from Landsat-5/7/8 and
207 Sentinel-2 during 1990–2024 and processed on the Google Earth Engine (GEE) cloud platform
208 (Table 1). Level-2 surface reflectance products were employed to ensure radiometric consistency



209 and accurate geolocation across sensors. To mitigate the effects of cloud contamination, images
210 with scene-level cloud cover exceeding 25% were excluded. Additional pixel-level masking was
211 applied using sensor-specific quality assessment (QA) bands to remove clouds, cloud shadows,
212 and cirrus contamination. For days with multiple valid observations, median compositing was
213 performed to generate daily surface reflectance mosaics with reduced cloud-related noise. Water
214 bodies were identified using the Modified Normalized Difference Water Index (MNDWI) (Xu,
215 2006), calculated as:

$$216 \quad \text{MNDWI} = \frac{\text{Green} - \text{SWIR}}{\text{Green} + \text{SWIR}} \quad (1)$$

217 Compared with the traditional Normalized Difference Water Index (NDWI) (McFeeters, 1996),
218 MNDWI more effectively suppresses signals from built-up land surfaces and improves the
219 delineation of open water bodies. Binary water masks were subsequently generated using the Otsu
220 adaptive thresholding algorithm (Otsu, 1979), which automatically determines the optimal
221 threshold by maximizing inter-class variance. Unlike fixed-threshold approaches, this adaptive
222 method accommodates spatiotemporal variability and has demonstrated robust performance across
223 diverse climatic, hydrological, and land-cover conditions (Chen and Zhao, 2022; Che et al., 2025).

224 **Table 1.** Summary of optical satellite datasets used in this study.

Data source	Sensor type	Time coverage	Spatial resolution (m)	Revisit period (days)	Spectral bands
Landsat-5	TM	1990–2011	30	16	Blue, Green, Red, NIR, SWIR1, Thermal IR, SWIR2
Landsat-7	ETM+	1999–2024	15, 30	16	Blue, Green, Red, NIR, SWIR1, SWIR2, Thermal IR, Panchromatic
Landsat-8	OLI	2013–2024	15, 30	16	Coastal aerosol, Blue, Green, Red, NIR, SWIR1, SWIR2, Panchromatic, Cirrus, Thermal IR1, Thermal IR2
Sentinel-2	MSI	2022–2024	10, 20, 60	5 (dual satellite)	Aerosol, Blue, Green, Red, Red Edge 1-4, NIR, Water Vapor, Cirrus, SWIR1, SWIR2

225 2.2.3 River width estimation and quality control

226 River width was measured along each predefined cross-section by sampling the binary water
227 mask at the native spatial resolution of each satellite sensor (30 m for Landsat and 10 m for
228 Sentinel-2). The instantaneous river width, W_t , was calculated as:

$$229 \quad W_t = L \times \frac{N_w}{N_{all}} \quad (2)$$



230 where L is the total cross-section length, defined as $3\bar{W}$; N_w is the number of sampling pixels
231 belonging to the largest connected water body along the cross-section; and N_{all} is the total number
232 of valid sampling points along the cross-section. To ensure the reliability of river width retrievals,
233 width estimates were retained only when N_{all} exceeded 80% of the theoretical maximum number
234 of sampling points (N_{max}). N_{max} is given by:

$$235 \quad N_{max} = \frac{L}{R} \quad (3)$$

236 where R denotes the spatial resolution of the satellite sensor (m). This quality-control criterion
237 effectively eliminates low-quality observations affected by cloud contamination, data gaps, or
238 geometric distortions. Additionally, river width estimates smaller than the sensor spatial resolution
239 were discarded as noise.

240 **2.3 River discharge reconstruction and validation**

241 The at-a-station hydraulic geometry (AHG), first introduced by Leopold and Maddock (1953),
242 describes the power-law relationship between river width and discharge at a specific cross-section:

$$243 \quad W = aQ^b \quad (4)$$

244 where a and b are empirically derived coefficients representing the scaling relationship between
245 river width and discharge. This relationship has been widely applied to characterize the nonlinear
246 response of river width to discharge variations and has demonstrated strong applicability for
247 satellite-based discharge estimation (Gleason and Wang, 2015; Riggs et al., 2023). Accordingly,
248 the instantaneous river widths derived from satellite imagery (Section 2.2) were paired with
249 concurrent in situ daily discharge observations at each gauge (Section 2.1). To ensure statistical
250 robustness, only gauges with at least ten valid width–discharge pairs ($N \geq 10$) were retained for
251 subsequent analysis.

252 Additional screening procedures were applied to remove gauges exhibiting weak or
253 physically unrealistic width–discharge relationships. First, the Spearman rank correlation
254 coefficient (ρ) was calculated to assess the monotonic association between river width and
255 discharge. Compared with linear correlation metrics, Spearman correlation is less sensitive to
256 outliers and nonlinear behavior and is therefore more suitable for hydraulic geometry analysis
257 (Patidar et al., 2025). Gauges with $\rho < 0.25$ were discarded. Such weak or negative correlations
258 generally indicate that river width does not reliably respond to discharge variability, often due to
259 strong channel confinement, backwater effects, reservoir regulation, or other anthropogenic



260 disturbances (Eggleston et al., 2024; Wang and Smith, 2025). Second, the fitted hydraulic
 261 geometry exponent b was evaluated for physical plausibility. In natural river systems, b typically
 262 ranges between 0 and 0.59, reflecting realistic channel width adjustments under varying flow
 263 conditions (Park, 1977; Yuan et al., 2024). Gauges with fitted exponents outside this range were
 264 removed to avoid incorporating unstable or physically implausible hydraulic relationships into the
 265 discharge reconstruction framework.

266 To quantitatively evaluate the performance of the satellite-based discharge reconstruction,
 267 four commonly used metrics were employed: the Pearson correlation coefficient (CC), normalized
 268 root mean square error (NRMSE), relative bias (rBias), and Kling–Gupta efficiency (KGE). These
 269 metrics assess different aspects of model performance, including correlation, error magnitude,
 270 systematic bias, and overall hydrological consistency. Their formulations are defined as follows:

$$271 \quad CC = \frac{\sum_{i=1}^N (Q_i^{\text{obs}} - \overline{Q^{\text{obs}}})(Q_i^{\text{est}} - \overline{Q^{\text{est}}})}{\sqrt{\sum_{i=1}^N (Q_i^{\text{obs}} - \overline{Q^{\text{obs}}})^2} \sqrt{\sum_{i=1}^N (Q_i^{\text{est}} - \overline{Q^{\text{est}}})^2}} \quad (5)$$

$$272 \quad NRMSE = \frac{\sqrt{\frac{1}{N} \sum_{i=1}^N (Q_i^{\text{est}} - Q_i^{\text{obs}})^2}}{Q_{\text{max}}^{\text{obs}} - Q_{\text{min}}^{\text{obs}}} \quad (6)$$

$$273 \quad rBias = \frac{\sum_{i=1}^N (Q_i^{\text{est}} - Q_i^{\text{obs}})}{\sum_{i=1}^N Q_i^{\text{obs}}} \quad (7)$$

274 where Q_i^{obs} and Q_i^{est} denote the observed and estimated discharge (m^3/s) at time step i ,
 275 respectively. $\overline{Q^{\text{obs}}}$ and $\overline{Q^{\text{est}}}$ represent the corresponding mean discharge values, and N is the total
 276 number of paired observations.

$$277 \quad KGE = 1 - \sqrt{(r - 1)^2 + (\alpha - 1)^2 + (\beta - 1)^2} \quad (8)$$

$$278 \quad r = CC, \alpha = \sigma^{\text{est}} / \sigma^{\text{obs}}, \beta = \mu^{\text{est}} / \mu^{\text{obs}} \quad (9)$$

279 where σ^{obs} and σ^{est} represent the standard deviations of the observed and estimated discharge,
 280 respectively, while μ^{obs} and μ^{est} denote their mean values.

281 Compared with traditional performance metrics such as the Nash–Sutcliffe efficiency (NSE),
 282 KGE explicitly decomposes model performance into correlation, variability, and bias components,
 283 thereby providing a more balanced and diagnostically informative evaluation of hydrological
 284 simulations. Owing to these advantages, KGE has been widely adopted in recent studies of
 285 remotely sensed discharge estimation, facilitating direct comparison with existing datasets and
 286 methodologies (Lin et al., 2023; Riggs et al., 2023; Elmi et al., 2024; Saemian et al., 2024).

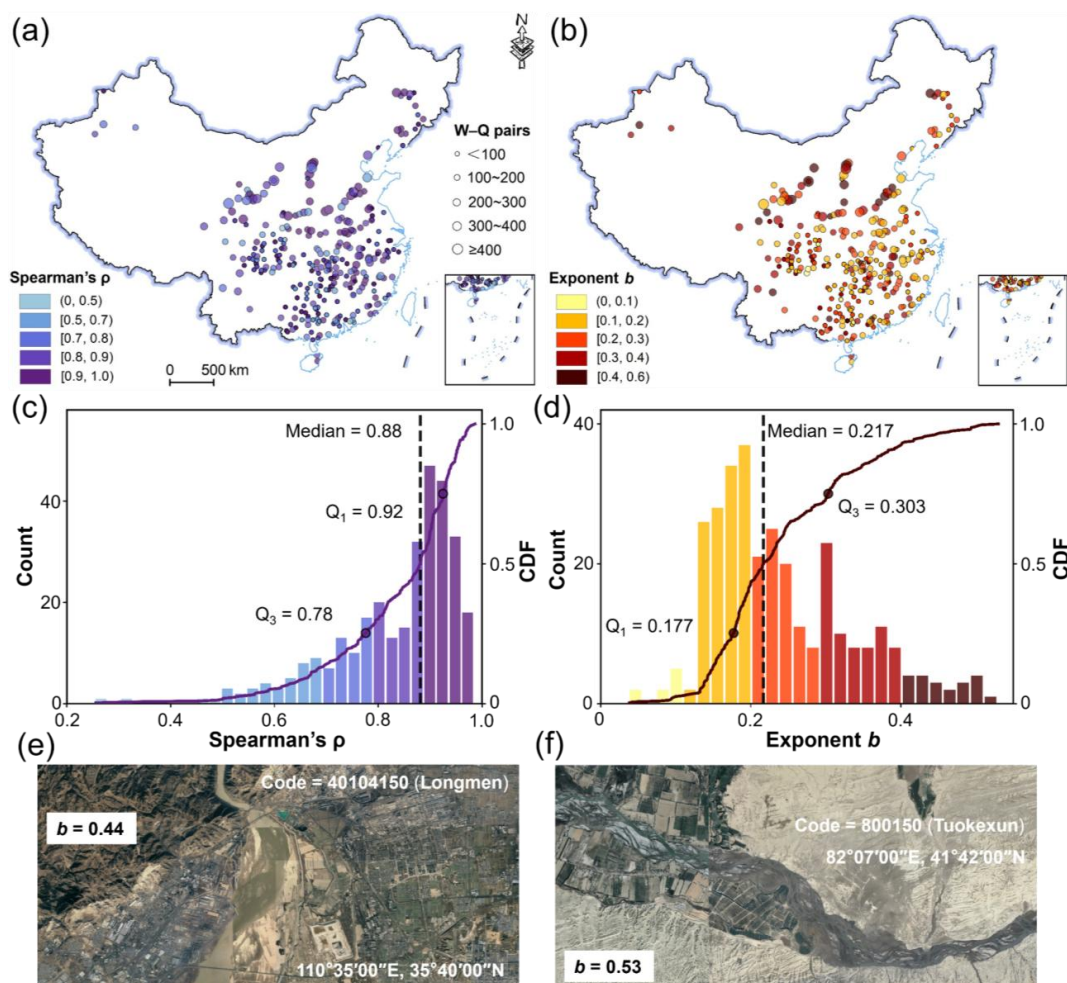


287 3 Results

288 3.1 Evaluation of width–discharge relationships

289 Following the screening criteria, a total of 310 gauges with sufficient width–discharge
290 observations and physically consistent hydraulic geometry relationships were retained for
291 discharge reconstruction (Fig. 5). The ρ between river width and discharge ranges from 0.26 to
292 0.99, with a median value of 0.88 and a mean of 0.84 (Fig. 5a, c). The interquartile range (IQR)
293 spans 0.78–0.92, and approximately 70% of the gauges exhibit strong correlations ($\rho > 0.8$),
294 whereas only 1.3% show correlations below 0.5. Additionally, nearly 90% of gauges fall within
295 the range of 0.59–0.96. Spatially, strong width–discharge correlations are widely distributed across
296 diverse hydroclimatic regions of China, whereas weak correlations show no clear geographic
297 clustering (Fig. 5a). These results indicate that river width at the selected gauges responds
298 sensitively and consistently to discharge variability, thereby satisfying a key prerequisite for
299 satellite-based discharge estimation.

300 The hydraulic geometry exponent b derived from the AHG relationship further supports the
301 physical realism of the width–discharge relationship (Fig. 5b, d). The fitted b values range from
302 0.039 to 0.530, with a median of 0.217, a mean of 0.243, and an IQR of 0.177–0.303.
303 Approximately 92% of gauges exhibit b values between 0 and 0.4, consistent with previous large-
304 sample hydraulic geometry studies. For example, Park (1977) reported a median b value of ~ 0.23
305 based on 139 gauges, while Yuan et al. (2024) found a median value of 0.213 using 436 global
306 gauges, with 94% of gauges falling within the 0–0.4 interval. Notably, relatively high b values
307 (e.g., 0.44 and 0.53) are primarily associated with gauges located in multi-threaded or laterally
308 unconfined river systems (Fig. 5e, f). In these environments, increasing discharge often induces
309 substantial lateral inundation across multiple channels, bars, and floodplain surfaces, thereby
310 enhancing the sensitivity of river width to discharge fluctuations. Such amplified width responses
311 are characteristics of braided and morphodynamically active river systems (Ashmore and Sauks,
312 2006; Yuan et al., 2024). Overall, these results demonstrate that satellite-derived river widths
313 effectively capture real discharge dynamics and provide a physically robust basis for extending
314 daily river discharge records across the selected gauges.

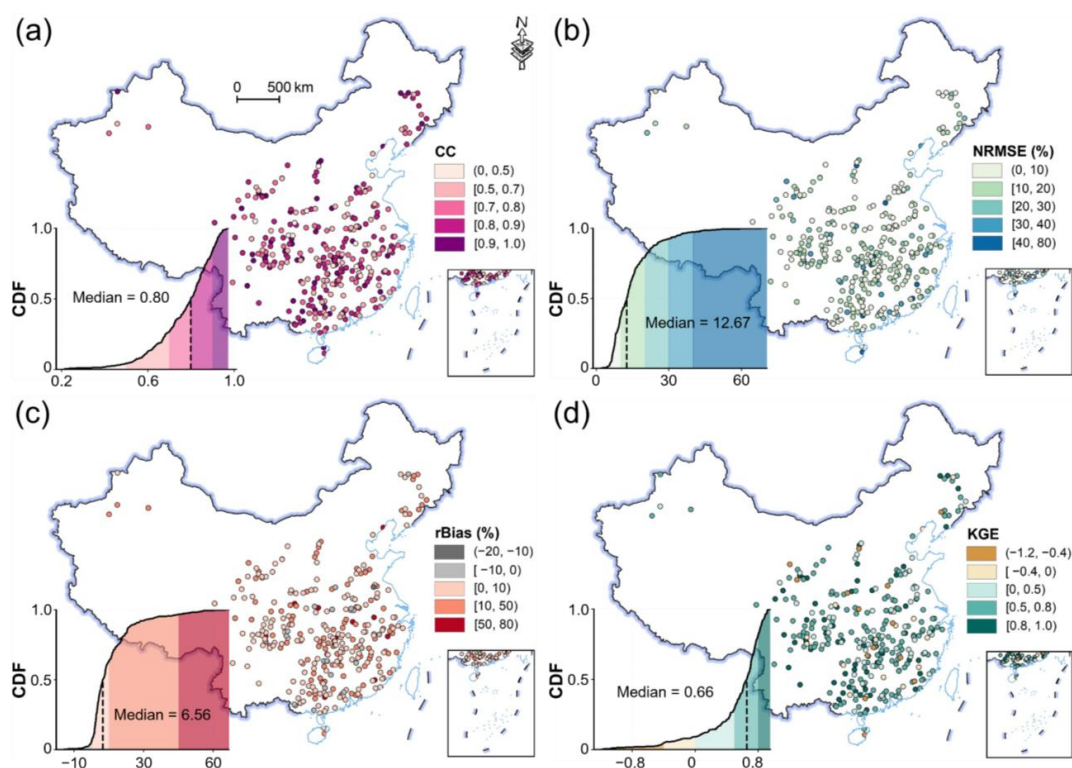


315
 316 **Figure 5.** Spatial distribution and statistical characteristics of width–discharge relationships across
 317 the 310 selected gauges. (a) Spatial distribution of the Spearman rank correlation coefficient (ρ)
 318 between river width and discharge, with symbol size representing the number of width–discharge
 319 (W–Q) observation pairs at each gauge. (b) Spatial distribution of the power-law exponent (b)
 320 derived from the at-a-station hydraulic geometry (AHG) relationship. (c–d) Frequency histograms
 321 and cumulative distribution functions (CDFs) of ρ and b , respectively, with median values and
 322 interquartile ranges indicated. (e–f) High-resolution satellite imagery of representative river
 323 reaches exhibiting high b values and morphologically dynamic channel patterns. Source of the
 324 underlying image: © Google Earth.



325 **3.2 Performance evaluation of satellite-extended river discharge**

326 Building upon the robust width–discharge power-law relationships established at 310 gauges,
327 we reconstructed and extended daily river discharge records for the period 1990–2024 and
328 evaluated performance using four complementary metrics. Overall, the satellite-extended
329 discharge exhibits strong agreement with in situ observations (Fig. 6). The median values of the
330 CC, NRMSE, rBias, and KGE are 0.80, 12.67%, 6.56%, and 0.66, respectively. Approximately
331 73% of gauges exhibit strong correlation ($CC > 0.7$), while 79% achieve NRMSE values below
332 20%, indicating that the reconstructed discharge reliably captures both the temporal variability and
333 magnitude of observations across diverse hydrological regimes. Notably, KGE values exceed
334 -0.41 at 96% of gauges, suggesting that the reconstructed discharge consistently outperforms a
335 simple mean discharge benchmark and retains meaningful predictive skill (Knoben et al., 2019).
336 Moreover, 91% of gauges show positive KGE values, demonstrating that the reconstructed records
337 generally reproduce observed discharge dynamics with acceptable overall fidelity. Spatially,
338 gauges with high predictive performance are broadly distributed across China’s major river basins,
339 with no obvious concentration in any specific hydroclimatic region. This widespread performance
340 underscores the robustness and scalability of the proposed framework and demonstrates the strong
341 potential of integrating satellite-derived river width with AHG relationships to reconstruct daily
342 river discharge over large spatial domains.

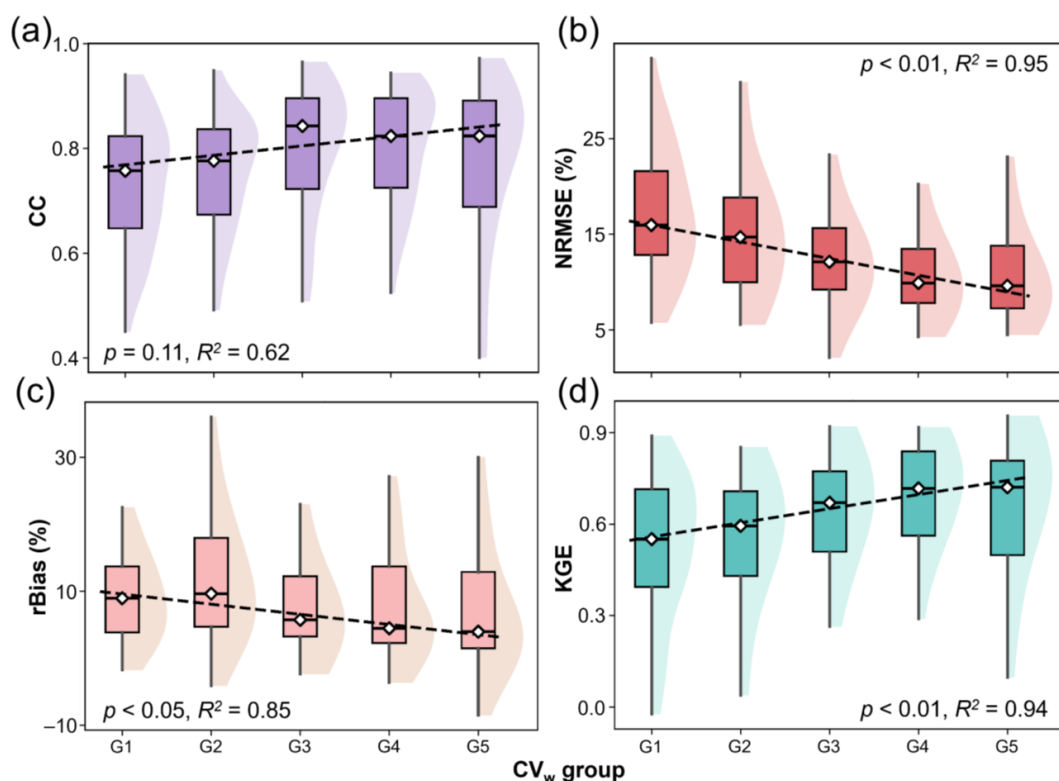


343
344 **Figure 6.** Performance evaluation of satellite-extended river discharge across 310 gauges based
345 on the AHG width–discharge power-law relationship. (a–d) Spatial distributions of the Pearson
346 correlation coefficient (CC), normalized root mean square error (NRMSE, %), relative bias
347 (rBias, %), and Kling–Gupta efficiency (KGE), respectively. Insets in each panel present the
348 corresponding cumulative distribution functions (CDFs), with median values indicated.

349 To further investigate how reconstruction performance depends on river dynamism, the
350 selected gauges were stratified into five equal-frequency groups (G1–G5) according to the
351 coefficient of variation of river width (CV_w), with group mean values ranging from 6% to 23%
352 (Fig. 7). Performance metrics were then compared across the CV_w gradient. The results reveal a
353 clear positive relationship between discharge reconstruction performance and river width
354 variability. Specifically, both NRMSE and rBias decrease with increasing CV_w (Fig. 7b, c), while
355 KGE increases correspondingly (Fig. 7d), with all trends statistically significant ($p < 0.05$). These
356 findings are physically consistent with expectations, as rivers exhibiting greater morphological
357 variability generally respond more sensitively to discharge fluctuations. Such enhanced hydraulic
358 responsiveness improves the detectability of width changes in satellite observations and



359 strengthens the robustness of width-based discharge estimation (Durand et al., 2023; Lin et al.,
360 2023; Scherer et al., 2024). Although CC does not exhibit a statistically significant monotonic
361 trend ($p = 0.11$), the relatively high coefficient of determination ($R^2 = 0.62$) still suggests a
362 moderate positive association with CV_w (Fig. 7a). However, CC values remain consistently high
363 (>0.75) across all groups, with only limited intergroup variability, indicating that the temporal
364 dynamics of river discharge are effectively captured regardless of river width variability. Notably,
365 even within the low- CV_w groups, where predictive performance is comparatively weaker, the
366 reconstruction framework retains satisfactory skill, with median CC, NRMSE, rBias, and KGE
367 values of approximately 0.76, 15.97%, 9.65%, and 0.55, respectively. These results demonstrate
368 that reliable discharge reconstruction can still be achieved in rivers with relatively limited width
369 variability, highlighting the robustness and broad applicability of the proposed framework across
370 diverse channel morphologies and hydrological conditions.



371
372 **Figure 7.** Performance evaluation of satellite-extended discharge across gradients of river width
373 variability, quantified by the coefficient of variation of river width (CV_w). (a–d) Boxplots of
374 performance metrics, including the Pearson correlation coefficient (CC), normalized root mean



375 square error (NRMSE), relative bias (rBias), and Kling–Gupta efficiency (KGE), for the 310
376 gauges grouped into five equal-frequency classes according to CV_w . Outliers are not shown for
377 clarity. Dashed lines indicate linear regression fits to the group median values, with corresponding
378 significance levels (p) and coefficients of determination (R^2) shown in each panel.

379 **3.3 Long-term trends and observation frequency of satellite-extended river discharge**

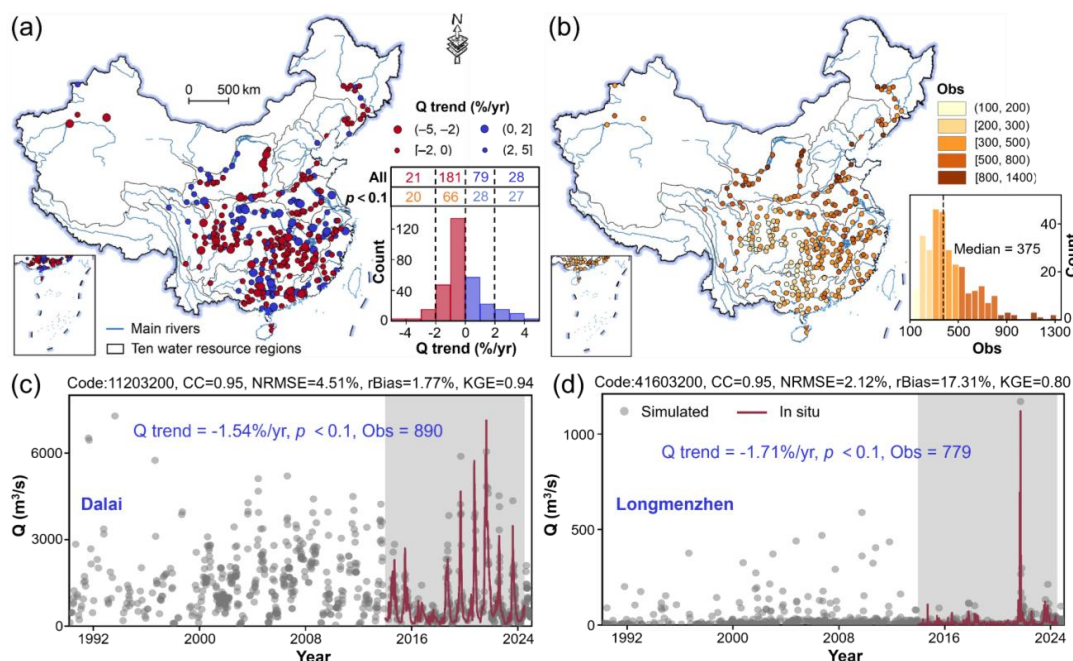
380 Using the satellite-extended daily discharge records from 310 gauges spanning 1990–2024,
381 we quantified the long-term trends and observational characteristics of rivers across China (Fig.
382 8). Trend detection was performed using the non-parametric Mann-Kendall test (Mann, 1945) in
383 combination with Sen’s slope estimator (Sen, 1968), applied to annually aggregated discharge
384 series. Relative annual rates of change (%/yr) were calculated by normalizing Sen’s slope with the
385 multi-year mean discharge, thereby facilitating comparison among rivers with varying discharge
386 magnitudes. To improve the robustness of trend estimation, only years containing at least three
387 observations separated by a minimum intra-annual interval of six months were retained.
388 Furthermore, trend analysis was restricted to gauges with more than 15 valid years of records.
389 After quality filtering, 309 gauges were included in the final analysis, with an average effective
390 record length of 32.5 years.

391 The results reveal a subtle but widespread decline in river discharge across China, with
392 approximately 65% ($n = 201$) of gauges exhibiting negative trends and a median relative change
393 rate of $-0.21\%/yr$ (Fig. 8a). Spatially, declining trends are concentrated in northern and central
394 China, particularly within the Haihe, Liaohe, and Yellow River basins, as well as parts of the
395 middle Yangtze River basin. This pattern is consistent with recent assessments suggesting that
396 discharge reductions in these regions are primarily driven by geomorphological alterations, further
397 intensified by anthropogenic water management activities such as water abstraction, inter-basin
398 diversion, and reservoir regulation (Yang et al., 2022; Wang et al., 2025a). In terms of trend
399 magnitude, most declining gauges ($n = 180$) exhibit relatively modest decreases of less than $2\%/yr$,
400 whereas 21 gauges show declines exceeding this threshold. Importantly, a persistent decrease
401 greater than $2\%/yr$ corresponds to a cumulative reduction exceeding 40% over the past three
402 decades, indicating substantial long-term discharge loss at these locations. Statistical significance
403 testing identifies monotonic trends at 141 gauges (46% of all analyzed gauges, $p < 0.1$), among
404 which 61% exhibit declining discharge. Within the significantly declining subset, 66 gauges



405 decrease by less than 2%/yr, while 20 gauges decline by more than 2%/yr. Conversely, 108 gauges
406 exhibit increasing discharge trends, of which approximately 51% are statistically significant. These
407 increasing trends are primarily concentrated in the northwestern and middle-to-lower Yangtze
408 River basin, the northern Pearl River basin, and the upper Yellow River basin. Overall, the spatial
409 heterogeneity of discharge trends reflects the diverse and complex hydrological responses
410 occurring across China's distinct geographical and hydroclimatic regions.

411 The temporal sampling characteristics of the satellite-extended discharge dataset were further
412 examined (Fig. 8b). The median number of observations per gauge is 375, corresponding to
413 approximately 11 observations per year, while about 42% of gauges contain more than 420
414 observations (≥ 12 observations per year). Despite this generally dense temporal coverage,
415 observation frequency exhibits substantial spatial heterogeneity across China. Gauges with fewer
416 than 300 observations are predominantly concentrated in the middle Yangtze and Pearl River
417 basins. This reduced sampling density is primarily associated with persistent cloud cover, frequent
418 precipitation, and complex topography, which limit the availability of high-quality optical satellite
419 imagery and consequently reduce the number of usable observations. Nevertheless, even under
420 irregular observational intervals, the reconstructed discharge records preserve key hydrological
421 dynamics. For example, the reconstructed time series at the Dalai (Fig. 8c) and Longmenzhen (Fig.
422 8d) gauges successfully capture transitions from baseflow to peak-flow conditions while
423 maintaining coherent multidecadal variability and long-term discharge trends. These examples
424 demonstrate that the satellite-based reconstruction framework can resolve both short-term
425 hydrological fluctuations and long-term discharge despite the discontinuous sampling inherent in
426 optical remote sensing observations. Overall, these results indicate that the satellite-extended
427 discharge dataset provides sufficient temporal depth and observational frequency to support robust
428 gauge-scale trend analysis, offering valuable insights into the long-term hydrological evolution of
429 China's major river systems.



430

431 **Figure 8.** Spatial patterns and long-term trends of satellite-extended river discharge across 309
 432 gauges in China (1990–2024). (a) Spatial distribution of relative discharge change rates (Q
 433 trend, %/yr), normalized by multi-year mean discharge, with inset histograms showing the trend
 434 distribution and corresponding significance levels (p) across all gauges. (b) Spatial distribution of
 435 additional observation (Obs) provided by the satellite-extended discharge records at each gauge.
 436 (c–d) Representative time series of reconstructed daily discharge at Dalai and Longmenzhen
 437 gauges, respectively, illustrating long-term discharge variability and trends together with
 438 corresponding model performance metrics.

439 4 Discussion

440 4.1 Comparison of reconstructed discharge records with existing datasets

441 To better contextualize the contribution of the CDR² dataset, we conducted a comparative
 442 assessment against several representative global remotely sensed river discharge products,
 443 including the Remote Sensing-based Extension for the GRDC (RSEG) (Elmi et al., 2024), the
 444 Satellite Altimetry-based Extension of global in situ river discharge Measurements (SAEM)
 445 (Saemian et al., 2024), and the datasets developed by Lin et al. (2023) and Riggs et al. (2023).
 446 Although these datasets have substantially expanded the spatial coverage of global river discharge



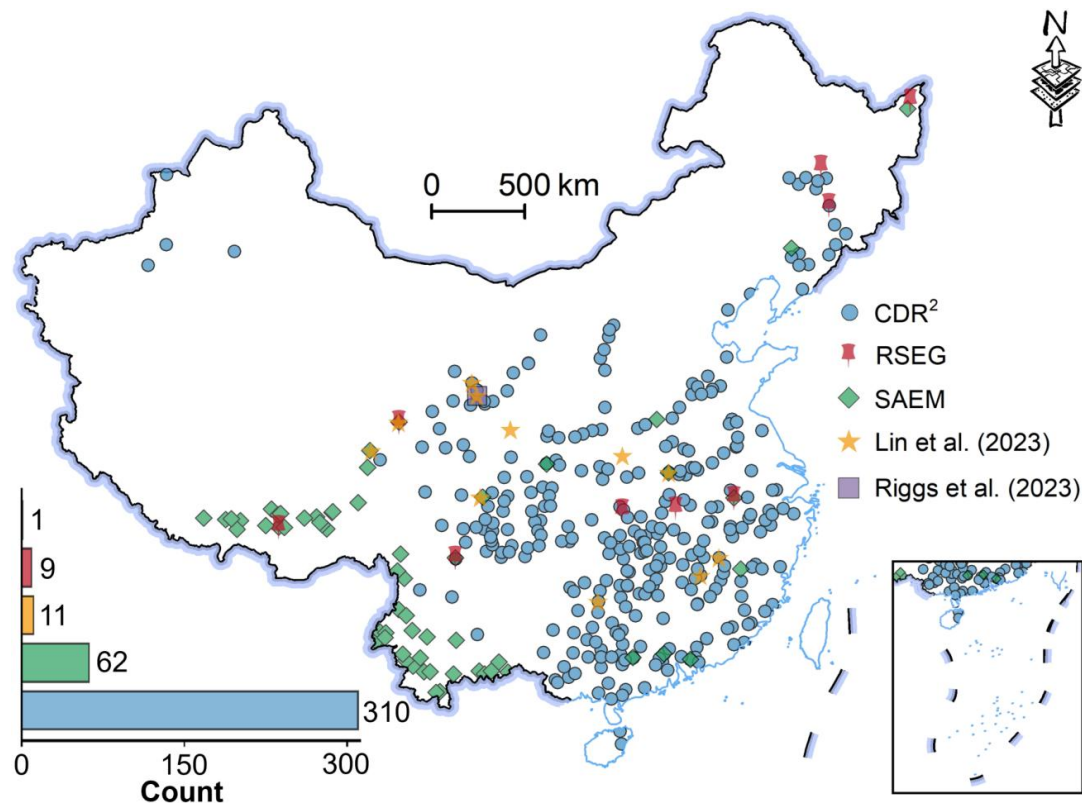
447 observations, their representation within China remains extremely limited (Fig. 9 and Table 2).
448 Specifically, the number of gauges in China ranges from only one in Riggs et al. (2023) to 62 in
449 SAEM. In all cases, Chinese gauges account for less than 1% of the total stations included in these
450 global products, reflecting both the scarcity of publicly accessible in situ discharge observations
451 and the challenges of satellite-based discharge retrieval in this region. In contrast, CDR² increases
452 the number of gauge-constrained discharge records within China by approximately 5–300 times
453 relative to existing global datasets, providing daily discharge reconstructions at 310 gauges
454 distributed across diverse hydroclimatic and geomorphological settings (Fig. 9). The dataset
455 encompasses rivers spanning a wide range of discharge magnitudes and channel sizes, thereby
456 substantially improving the representation of China within global river discharge archives.

457 Beyond its improved spatial coverage, CDR² also demonstrates clear advantages in temporal
458 resolution and predictive performance. While RSEG and Lin et al. (2023) provide monthly
459 discharge estimates, CDR², SAEM, and Riggs et al. (2023) operate at daily time scales. Notably,
460 CDR² achieves a median KGE of 0.66, substantially outperforming the reported global median
461 range of 0.33–0.48 among existing products (Table 2). Consistent improvements are also evident
462 in other performance metrics, including a lower median NRMSE (12.7%) and a higher median CC
463 (0.80), indicating enhanced capability in reproducing both the magnitude and temporal variability
464 of observed discharge.

465 These improvements are partly attributed to the rigorous screening strategy adopted in this
466 study, whereby only gauges exhibiting stable width–discharge relationships and physically
467 consistent hydraulic behavior were retained (Fig. 5). In contrast, global-scale products often face
468 an inherent trade-off between maximizing spatial coverage and maintaining retrieval quality,
469 frequently incorporating gauges with weaker hydraulic relationships or unstable channel responses
470 that introduce additional uncertainty. Moreover, by integrating observations from both Landsat
471 and Sentinel-2 imagery, CDR² achieves a substantially higher effective sampling frequency than
472 single-sensor approaches. This enhanced temporal sampling improves detection of short-term
473 discharge fluctuations and mitigates the smoothing effects of monthly aggregation, as observed in
474 RSEG and Lin et al. (2023), thereby enabling a more realistic reconstruction of daily discharge
475 dynamics. Overall, CDR² provides a substantially improved and higher-resolution characterization
476 of river discharge across China, a hydrologically critical yet globally underrepresented region.



477 Through its greatly expanded gauge density, daily temporal resolution, and enhanced predictive
 478 performance, the dataset fills a major geographic gap in existing global river discharge products.



479
 480 **Figure 9.** Spatial distribution of river discharge gauges included in CDR², compared with gauges
 481 from RSEG (Elmi et al., 2024), SAEM (Saemian et al., 2024), Lin et al. (2023), and Riggs et al.
 482 (2023). The inset summarizes the number of gauges contained in each dataset within China.

483 **Table 2.** Comparison of CDR² with existing satellite-extended river discharge datasets.

Dataset	Hydraulic variables	Data source	Global gauge	China gauge	Temporal resolution	Validation results
Elmi et al., 2024 (RSEG)	River width/stage	Landsat/Satellite altimetry	3377	9	Monthly	All gauges (mean): KGE = 0.35, rRMSE = 21.8%, CC = 0.43
Saemian et al., 2024 (SAEM)	River stage	Satellite altimetry	8730	62	Daily	All gauges (median): KGE = 0.48, nRMSE = 18%, CC = 0.64
Lin et al., 2023	River width	Landsat	3078	11	Monthly	>600 gauges (median): KGE = 0.33
Riggs et al., 2023	River width	Landsat/Sentinel	2168	1	Daily	All gauges (median): KGE = 0.46, rRMSE = 83%, rBias = 1.4%
CDR ²	River width	Landsat/Sentinel	310	310	Daily	All gauges (median): KGE = 0.66, nRMSE = 12.7%, rBias = 6.6%, CC = 0.8



484 4.2 Sensitivity of reconstruction performance to the width–discharge exponent

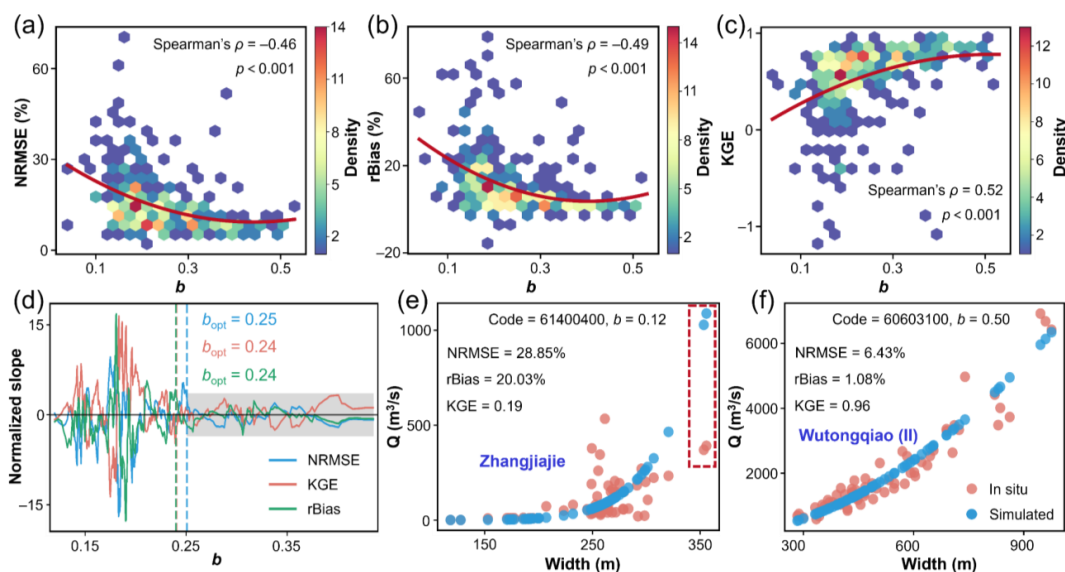
485 To investigate how hydraulic controls influence discharge reconstruction accuracy, we
486 evaluated the sensitivity of model performance to the width–discharge exponent b (Fig. 10). This
487 exponent characterizes the responsiveness of river width to discharge variability and therefore
488 serves as a key hydraulic link between satellite-observed channel dynamics and river discharge
489 behavior. The analysis reveals statistically significant relationships between b and multiple
490 performance metrics (Fig. 10a–c). Specifically, both NRMSE and rBias exhibit significant
491 negative correlations with b ($\rho = -0.46$ and -0.49 , respectively; $p < 0.001$), whereas KGE shows
492 a significantly positive correlation ($\rho = 0.52$, $p < 0.001$). These results indicate that rivers with
493 greater hydraulic sensitivity experience larger width variations under comparable discharge
494 fluctuations, making these changes more readily detectable by satellite observations (Fig. 7) and
495 thereby mitigating the propagation of observational uncertainty into discharge reconstruction.

496 To identify a robust threshold for hydraulic sensitivity, gauges were first sorted according to
497 increasing b values, after which a sliding-window analysis was applied to evaluate local gradients
498 of each performance metric with respect to b . To facilitate comparison across metrics with different
499 numerical ranges, the gradients were normalized to their respective global ranges. The optimal
500 threshold (b_{opt}) was defined as the point at which the absolute normalized gradient consistently
501 declined below 20% of its global maximum, indicating that performance stabilizes beyond this
502 value. The intersection of these stabilization regions across all metrics yields a global threshold of
503 $b = 0.25$ (Fig. 10d). Below this threshold, large fluctuations in normalized gradients indicate
504 considerable uncertainty in discharge reconstruction for hydraulically constrained rivers. Once b
505 exceeds 0.25, all gradients converge toward zero, suggesting that width-based discharge estimation
506 becomes substantially more stable and reliable for rivers with stronger hydraulic responsiveness.

507 The contrasting influence of b on reconstruction performance is clearly illustrated by the
508 Zhangjiajie and Wutongqiao (II) gauges (Fig. 10e, f). At Zhangjiajie, characterized by a low b
509 value, substantial discharge variability corresponds to only marginal changes in river width, often
510 approaching the spatial resolution limit of the satellite sensor (~ 30 m). Under such conditions, the
511 inversion from width to discharge becomes mathematically ill-conditioned, whereby small
512 uncertainties in width retrieval propagate into disproportionately large discharge errors, ultimately
513 degrading reconstruction performance (Fig. 10e). In contrast, the Wutongqiao (II) gauge exhibits
514 a high b value, where pronounced width sensitivity to discharge allows satellite observations to



515 effectively resolve discharge variations across the full hydrological spectrum, from low-flow
 516 conditions to flood peaks. Together, these findings demonstrate that b provides a physically
 517 meaningful and robust criterion for identifying rivers suitable for reliable satellite-based discharge
 518 reconstruction, offering important guidance for future large-scale discharge estimation and
 519 monitoring efforts.



520
 521 **Figure 10.** Influence of the width–discharge power-law exponent (b) on the performance of
 522 satellite-extended discharge reconstruction. (a–c) Relationships between the hydraulic geometry
 523 exponent (b) and the performance metrics NRMSE, rBias, and KGE, respectively. Red solid lines
 524 indicate fitted relationships, with the corresponding Spearman rank correlation coefficient (ρ) and
 525 significance level (p) indicated. (d) Variation in the normalized slope of performance metrics with
 526 respect to b , used to identify the optimal threshold (b_{opt}) beyond which reconstruction performance
 527 becomes stable and less sensitive to changes in hydraulic responsiveness. (e–f) Comparison
 528 between in situ and reconstructed discharge at the Zhangjiajie and Wutongqiao (II) gauges,
 529 representing examples of low- b and high- b hydraulic conditions, respectively.

530 5 Data availability

531 The China Daily River Discharge Records (CDR²) dataset developed in this study is publicly
 532 available on Zenodo at <https://doi.org/10.5281/zenodo.20152832> (Wang and Li, 2026). The
 533 dataset is distributed in CSV format, with each file named according to its corresponding gauge



534 ID. It contains reconstructed daily river discharge records for 310 gauges across China, along with
535 the associated width–discharge power-law relationships, covering the period from 1990 to 2024.

536 The in situ daily river discharge observations were sourced from the National Hydrological
537 and Rainfall Information System of the Ministry of Water Resources of China
538 (<http://xxfb.mwr.cn/>). The Global Runoff Data Centre (GRDC) data can be accessed
539 at:<https://portal.grdc.bafg.de/applications/public.html?publicuser=PublicUser>. The Chinese
540 Hydrology Project (CHP) data are not publicly available and were provided by the dataset’s
541 authors (Henck et al., 2010; Schmidt et al., 2011). The Surface Water and Ocean Topography
542 River Database (Version v17) is available at: <https://zenodo.org/records/14727521>. All Landsat-
543 5/7/8 and Sentinel-2 optical remote sensing data were accessed via the Google Earth Engine
544 platform (<https://earthengine.google.com/>). The following publicly available global satellite-based
545 river discharge products were used for comparison. Satellite Altimetry-based Extension of global
546 in situ river discharge Measurements (SAEM) is available at: <https://doi.org/10.18419/darus-4475>.
547 Remote Sensing-based Extension for the GRDC (RSEG) is available at:
548 <https://doi.org/10.18419/darus-3558>. Datasets developed by Lin et al. (2023) and Riggs et al.
549 (2023), which can be accessed at: <https://zenodo.org/records/6655532#.Yqw37XZBz-g> and
550 <https://zenodo.org/record/7150168#.ZAow8RXMIuV>.

551 **6 Conclusions**

552 This study integrates multi-decadal optical remote sensing observations, a robust at-a-station
553 hydraulic geometry (AHG) width–discharge framework, and in situ daily discharge records from
554 1,196 gauging stations to develop the China Daily River Discharge Records (CDR²), a high-
555 resolution satellite-extended daily discharge dataset spanning 1990–2024 for 310 gauges across
556 China. Validation against in situ observations demonstrates that CDR² accurately captures both
557 the temporal dynamics and magnitude of river discharge, achieving median values of 0.66 for
558 Kling–Gupta Efficiency (KGE), 0.80 for the Pearson correlation coefficient (CC), 6.56% for
559 relative bias (rBias), and 12.7% for normalized root mean square error (NRMSE). Notably, 91%
560 of gauges exhibit positive KGE values, approximately 73% achieve CC values greater than 0.7,
561 and 79% maintain NRMSE below 20%.

562 Compared with existing global satellite-based discharge datasets, CDR² expands the spatial
563 coverage of daily river discharge monitoring across China by more than fivefold and demonstrates



564 substantially improved accuracy. Sensitivity analyses further reveal a strong positive relationship
565 between reconstruction performance and river width variability, indicating that satellite
566 observations more effectively resolve rivers exhibiting greater hydraulic responsiveness to
567 discharge fluctuations. Long-term trend analysis reveals a widespread decline in river discharge
568 across China over the past 35 years, with approximately 65% of gauges exhibiting decreasing
569 trends. The median relative rate of change is $-0.21\%/yr$, corresponding to an estimated cumulative
570 reduction of $\sim 7\%$ relative to the long-term mean discharge. These decreases are particularly
571 pronounced in northern and central basins, highlighting increasing spatial disparities in regional
572 hydrological dynamics and underscoring the need for sustained, high-quality discharge monitoring.
573 By substantially expanding gauge-scale discharge observations across China, CDR² provides a
574 critical empirical foundation for calibrating next-generation satellite missions and advancing large-
575 scale river discharge estimates in the era of the Surface Water and Ocean Topography (SWOT).

576 **Author contributions**

577 YW performed the investigation, methodology, formal analysis, software, data curation,
578 visualization, and validation, and wrote the original draft. YG curated the data and reviewed and
579 edited the manuscript. YJ performed formal analysis and reviewed and edited the manuscript. XS
580 reviewed and edited the manuscript. YL conceptualized the study, developed the methodology,
581 conducted formal analysis, provided resources, supervised the research, administered the project,
582 acquired funding, and reviewed and edited the manuscript.

583 **Declaration of Competing Interest**

584 The authors declare that they have no known competing financial interests or personal
585 relationships that could have appeared to influence the work reported in this paper.

586 **Acknowledgments**

587 The authors gratefully acknowledge the data providers whose contributions were essential to this
588 study.

589 **Financial support**

590 This work was supported by the National Natural Science Foundation of China (42571446 and
591 42201349), the Chongqing Doctoral Student Research Innovation Project (CYB240107), and the



592 Chongqing Municipal Science and Technology Bureau (CSTB2024YCJH-KYXM0054 and
593 cstc2024ycjh-bgzxm0043).



594 References

- 595 Allen, G. H., Pavelsky, T. M.: Global extent of rivers and streams, *Science*, 361, 585–588,
596 <https://doi.org/10.1126/science.aat0636>, 2018.
- 597 Alsdorf, D. E., Rodríguez, E., Lettenmaier, D. P.: Measuring surface water from space, *Rev.*
598 *Geophys.*, 45, RG2002, <https://doi.org/10.1029/2006RG000197>, 2007.
- 599 Altenau, E. H., Pavelsky, T. M., Durand, M. T., Yang, X., Frasson, R. P. M., and Bendezu, L.: The
600 Surface Water and Ocean Topography (SWOT) Mission River Database (SWORD): A global
601 river network for satellite data products, *Water Resour. Res.*, 57, e2021WR030054,
602 <https://doi.org/10.1029/2021WR030054>, 2021.
- 603 Andrews, L. and Grantham, T. E.: Strategic stream gauging network design for sustainable water
604 management, *Nat. Sustain.*, 7, 714–723, <https://doi.org/10.1038/s41893-024-01357-z>, 2024.
- 605 Ashmore, P. and Sauks, E.: Prediction of discharge from water surface width in a braided river
606 with implications for at-a-station hydraulic geometry, *Water Resour. Res.*, 42, W03407,
607 <https://doi.org/10.1029/2005WR003993>, 2006.
- 608 Beven, K.: Towards integrated environmental models of everywhere: uncertainty, data and
609 modelling as a learning process, *Hydrol. Earth Syst. Sci.*, 11, 460–467,
610 <https://doi.org/10.5194/hess-11-460-2007>, 2007.
- 611 Bonnema, M. G., Sikder, S., Hossain, F., Durand, M., Gleason, C. J., and Bjerklie, D. M.:
612 Benchmarking wide swath altimetry-based river discharge estimation algorithms for the
613 Ganges river system, *Water Resour. Res.*, 52, 2439–2461,
614 <https://doi.org/10.1002/2015WR017296>, 2016.
- 615 Che, L., Li, S., and Liu, X.: Improved surface water mapping using satellite remote sensing
616 imagery based on optimization of the Otsu threshold and effective selection of remote-sensing
617 water index, *J. Hydrol.*, 654, 132771, <https://doi.org/10.1016/j.jhydrol.2025.132771>, 2025.
- 618 Chen, Z. and Zhao, S.: Automatic monitoring of surface water dynamics using Sentinel-1 and
619 Sentinel-2 data with Google Earth Engine, *Int. J. Appl. Earth Obs. Geoinf.*, 113, 103010,
620 <https://doi.org/10.1016/j.jag.2022.103010>, 2022.
- 621 Collins, E. L., David, C. H., Riggs, R., Allen, G. H., Pavelsky, T. M., Lin, P., Pan, M., Yamazaki,
622 D., Meentemeyer, R. K., and Sanchez, G. M.: Global patterns in river water storage dependent
623 on residence time, *Nat. Geosci.*, 17, 433–439, <https://doi.org/10.1038/s41561-024-01421-5>,
624 2024.



- 625 Do, H. X., Gudmundsson, L., Leonard, M., and Westra, S.: The Global Streamflow Indices and
626 Metadata Archive (GSIM)–Part 1: The production of a daily streamflow archive and metadata,
627 Earth Syst. Sci. Data, 10, 765–785, <https://doi.org/10.5194/essd-10-765-2018>, 2018.
- 628 Durand, M., Gleason, C. J., Garambois, P. A., Bjerklie, D., Smith, L. C., Roux, H., Rodriguez, E.,
629 Bates, P. D., Pavelsky, T. M., Monnier, J., Chen, X., Di Baldassarre, G., Fiset, J.-M., Flipo,
630 N., Frasson, R. P. d. M., Fulton, J., Goutal, N., Hossain, F., Humphries, E., Minear, J. T.,
631 Mukolwe, M. M., Neal, J. C., Ricci, S., Sanders, B. F., Schumann, G., Schubert, J. E., and
632 Vilmin, L.: An intercomparison of remote sensing river discharge estimation algorithms from
633 measurements of river height, width, and slope, Water Resour. Res., 52, 4527–4549,
634 <https://doi.org/10.1002/2015WR018434>, 2016.
- 635 Durand, M., Gleason, C. J., Pavelsky, T. M., Frasson, R. P. M., Turmon, M., David, C. H., Altenau,
636 E. H., Tebaldi, N., Larnier, K., Monnier, J., Malaterre, P. O., Oubanas, H., Allen, G. H.,
637 Astifan, B., Brinkerhoff, C., Bates, P. D., Bjerklie, D., Coss, S., Dudley, R., Fenoglio, L.,
638 Garambois, P.-A., Getirana, A., Lin, P., Margulis, S. A., Matte, P., Minear, J. T., Muhebwa,
639 A., Pan, M., Peters, D., Riggs, R., Sikder, M. S., Simmons, T., Stuurman, C., Taneja, J.,
640 Tarpanelli, A., Schulze, K., Tourian, M. J., and Wang, J.: A framework for estimating global
641 river discharge from the Surface Water and Ocean Topography satellite mission, Water
642 Resour. Res., 59, e2021WR031614, <https://doi.org/10.1029/2021WR031614>, 2023.
- 643 Eggleston, J., Mason, C., Bjerklie, D., Durand, M., Dudley, R., and Harlan, M.: Siting
644 considerations for satellite observation of river discharge, Water Resour. Res., 60,
645 e2023WR034583, <https://doi.org/10.1029/2023WR034583>, 2024.
- 646 Elmi, O., Tourian, M. J., Saemian, P., and Sneeuw, N.: Remote sensing-based extension of GRDC
647 Discharge time series – a monthly product with uncertainty estimates, Sci. Data, 11, 240,
648 <https://doi.org/10.1038/s41597-024-03078-6>, 2024.
- 649 Famiglietti, J. S., Cazenave, A., Eicker, A., Reager, J. T., Rodell, M., and Velicogna, I.: Satellites
650 provide the big picture, Science, 349, 684–685, <https://doi.org/10.1126/science.aac9238>,
651 2015.
- 652 Färber, C., Plessow, H., Mischel, S. A., Kratzert, F., Addor, N., Shalev, G., and Looser, U.:
653 GRDC-Caravan: extending caravan with data from the global runoff data centre, Earth Syst.
654 Sci. Data, 17, 4613–4625, <https://doi.org/10.5194/essd-17-4613-2025>, 2025.



- 655 Fekete, B. M., Roberts, R. D., Kumagai, M., Nachtnebel, H.-P., Odada, E., and Zhulidov, A. V.:
656 Time for in situ renaissance, *Science*, 349, 685–686, <https://doi.org/10.1126/science.aac7358>,
657 2015.
- 658 Feng, D., Gleason, C. J., Yang, X., and Pavelsky, T. M.: Comparing discharge estimates made via
659 the BAM algorithm in high-order Arctic rivers derived solely from optical CubeSat, Landsat,
660 and Sentinel-2 data, *Water Resour. Res.*, 55, 7753–7771,
661 <https://doi.org/10.1029/2019WR025599>, 2019.
- 662 Ghiggi, G., Humphrey, V., Seneviratne, S. I., and Gudmundsson, L.: GRUN: an observation-based
663 global gridded runoff dataset from 1902 to 2014, *Earth Syst. Sci. Data*, 11, 1655–1674,
664 <https://doi.org/10.5194/essd-11-1655-2019>, 2019.
- 665 Gleason, C. J. and Smith, L. C.: Toward global mapping of river discharge using satellite images
666 and at-many-stations hydraulic geometry, *Proc. Natl. Acad. Sci. U. S. A.*, 111, 4788–4791,
667 <https://doi.org/10.1073/pnas.1317606111>, 2014.
- 668 Gleason, C. J. and Wang, J.: Theoretical basis for at-many-stations hydraulic geometry, *Geophys.*
669 *Res. Lett.*, 42, 7107–7114, <https://doi.org/10.1002/2015GL064935>, 2015.
- 670 Hannah, D. M., Demuth, S., van Lanen, H. A. J., Looser, U., Prudhomme, C., Rees, G., Stahl, K.,
671 and Tallaksen, L. M.: Large-scale river flow archives: Importance, current status and future
672 needs, *Hydrol. Process.*, 25, 1191–1200, <https://doi.org/10.1002/hyp.7794>, 2011.
- 673 Hao, Z., Xiang, N., Cai, X., Zhong, M., Jin, J., Du, Y., and Ling, F.: Remote sensing of river
674 discharge from medium-resolution satellite imagery based on deep learning, *Water Resour.*
675 *Res.*, 60, e2023WR036880, <https://doi.org/10.1029/2023WR036880>, 2024.
- 676 Henck, A. C., Montgomery, D. R., Huntington, K. W., and Liang, C.: Monsoon control of effective
677 discharge, Yunnan and Tibet, *Geology*, 38, 975–978, <https://doi.org/10.1130/G31444.1>, 2010.
- 678 Hou, J., van Dijk, A. I. J. M., and Beck, H. E.: Global satellite-based river gauging and the
679 influence of river morphology on its application, *Remote Sens. Environ.*, 239, 111629,
680 <https://doi.org/10.1016/j.rse.2019.111629>, 2020.
- 681 Knoben, W. J. M., Freer, J. E., and Woods, R. A.: Inherent benchmark or not? Comparing Nash–
682 Sutcliffe and Kling–Gupta efficiency scores, *Hydrol. Earth Syst. Sci.*, 23, 4323–4331,
683 <https://doi.org/10.5194/hess-23-4323-2019>, 2019.
- 684 Krabbenhoft, C. A., Allen, G. H., Lin, P., Godsey, S. E., Allen, D. C., Burrows, R. M., DelVecchia,
685 A. G., Fritz, K. M., Shanafield, M., Burgin, A. J., Zimmer, M. A., Datry, T., Dodds, W. K.,



- 686 Jones, C. N., Mims, M. C., Franklin, C., Hammond, J. C., Zipper, S., Ward, A. S., Costigan,
687 K. H., Beck, H. E., and Olden, J. D.: Assessing placement bias of the global river gauge
688 network, *Nat. Sustain.*, 5, 586–592, <https://doi.org/10.1038/s41893-022-00873-0>, 2022.
- 689 Leopold, L. B. and Maddock, T.: The hydraulic geometry of stream channels and some
690 physiographic implications, U.S. Government Printing Office, Washington, DC,
691 <https://pubs.usgs.gov/pp/0252/report.pdf> (last access: 12 March 2026), 1953.
- 692 Lettenmaier, D. P., Alsdorf, D., Dozier, J., Huffman, G. J., Pan, M., and Wood, E. F.: Inroads of
693 remote sensing into hydrologic science during the WRR era, *Water Resour. Res.*, 51, 7309–
694 7342, <https://doi.org/10.1002/2015WR017616>, 2015.
- 695 Lin, P., Pan, M., Beck, H. E., Yang, Y., Yamazaki, D., Frasson, R., David, C. H., Durand, M.,
696 Pavelsky, T. M., Allen, G. H., Gleason, C. J., and Wood, E. F.: Global reconstruction of
697 naturalized river flows at 2.94 million reaches, *Water Resour. Res.*, 55, 6499–6516,
698 <https://doi.org/10.1029/2019WR025287>, 2019.
- 699 Lin, P., Feng, D., Gleason, C. J., Pan, M., Brinkerhoff, C. B., Yang, X., Beck, H. E., and Frasson,
700 R. P. d. M.: Inversion of river discharge from remotely sensed river widths: A critical
701 assessment at three-thousand global river gauges, *Remote Sens. Environ.*, 287, 113489,
702 <https://doi.org/10.1016/j.rse.2023.113489>, 2023.
- 703 Mann, H. B.: Nonparametric tests against trend, *Econometrica*, 13, 245–259,
704 <https://doi.org/10.2307/1907187>, 1945.
- 705 McFeeters, S. K.: The use of the Normalized Difference Water Index (NDWI) in the delineation
706 of open water features, *Int. J. Remote Sens.*, 17, 1425–1432,
707 <https://doi.org/10.1080/01431169608948714>, 1996.
- 708 Oh, J. and Bartos, M.: Scalable, adaptive and risk-informed design of hydrological sensor networks,
709 *Nat. Water*, 3, 1144–1154, <https://doi.org/10.1038/s44221-025-00496-7>, 2025.
- 710 Oki, T. and Kanae, S.: Global hydrological cycles and world water resources, *Science*, 313, 1068–
711 1072, <https://doi.org/10.1126/science.1128845>, 2006.
- 712 Otsu, N.: A threshold selection method from gray-level histograms, *IEEE Trans. Syst. Man
713 Cybern.*, 9, 62–66, <https://doi.org/10.1109/TSMC.1979.4310076>, 1979.
- 714 Palmer, M. and Ruhi, A.: Measuring Earth's rivers, *Science*, 361, 546–547,
715 <https://doi.org/10.1126/science.aau3842>, 2018.



- 716 Paris, A., Dias de Paiva, R., Santos da Silva, J., Medeiros Moreira, D., Calmant, S., Garambois,
717 P.-A., Collischonn, W., Bonnet, M.-P., and Seyler, F.: Stage-discharge rating curves based
718 on satellite altimetry and modeled discharge in the Amazon basin, *Water Resour. Res.*, 52,
719 3787–3814, <https://doi.org/10.1002/2014WR016618>, 2016.
- 720 Park, C. C.: World-wide variations in hydraulic geometry exponents of stream channels: An
721 analysis and some observations, *J. Hydrol.*, 33, 133–146, [https://doi.org/10.1016/0022-
722 1694\(77\)90103-2](https://doi.org/10.1016/0022-1694(77)90103-2), 1977.
- 723 Patidar, G., Indu, J., and Karmakar, S.: Performance assessment of Surface Water and Ocean
724 Topography (SWOT) mission for WSE measurement across India, *Geophys. Res. Lett.*, 52,
725 e2025GL115804, <https://doi.org/10.1029/2025GL115804>, 2025.
- 726 Pavelsky, T. M. and Smith, L. C.: RivWidth: A software tool for the calculation of river widths
727 from remotely sensed imagery, *IEEE Geosci. Remote Sens. Lett.*, 5, 70–73,
728 <https://doi.org/10.1109/LGRS.2007.908305>, 2008.
- 729 Pavelsky, T. M.: Using width-based rating curves from spatially discontinuous satellite imagery
730 to monitor river discharge, *Hydrol. Process.*, 28, 3035–3040,
731 <https://doi.org/10.1002/hyp.10157>, 2014.
- 732 Resplandy, L., Keeling, R. F., Rödenbeck, C., Stephens, B. B., Khatiwala, S., Rodgers, K. B.,
733 Long, M. C., Bopp, L., and Tans, P. P.: Revision of global carbon fluxes based on a
734 reassessment of oceanic and riverine carbon transport, *Nat. Geosci.*, 11, 504–509,
735 <https://doi.org/10.1038/s41561-018-0151-3>, 2018.
- 736 Riggs, R. M., Allen, G. H., Wang, J., Pavelsky, T. M., Gleason, C. J., David, C. H., and Durand,
737 M.: Extending global river gauge records using satellite observations, *Environ. Res. Lett.*, 18,
738 064027, <https://doi.org/10.1088/1748-9326/acd407>, 2023.
- 739 Ruhi, A., Messenger, M. L., and Olden, J. D.: Tracking the pulse of the Earth's fresh waters, *Nat.*
740 *Sustain.*, 1, 198–203, <https://doi.org/10.1038/s41893-018-0047-7>, 2018.
- 741 Saemian, P., Elmi, O., Stroud, M., Riggs, R., Kitambo, B. M., Papa, F., Allen, G. H., and Tourian,
742 M. J.: Satellite Altimetry-based Extension of global-scale in situ river discharge
743 Measurements (SAEM), *Earth Syst. Sci. Data*, 17, 2063–2081, [https://doi.org/10.5194/essd-
744 17-2063-2025](https://doi.org/10.5194/essd-17-2063-2025), 2025.



- 745 Scherer, D., Schwatke, C., Dettmering, D., and Seitz, F.: Monitoring river discharge from space:
746 An optimization approach with uncertainty quantification for small ungauged rivers, *Remote*
747 *Sens. Environ.*, 315, 114434, <https://doi.org/10.1016/j.rse.2024.114434>, 2024.
- 748 Schmidt, A. H., Montgomery, D. R., Huntington, K. W., and Liang, C.: The question of communist
749 land degradation: New evidence from local erosion and basin-wide sediment yield in
750 Southwest China and Southeast Tibet, *Ann. Assoc. Am. Geogr.*, 101, 477–496,
751 <https://doi.org/10.1080/00045608.2011.560059>, 2011.
- 752 Sen, P. K.: Estimates of the regression coefficient based on Kendall's tau, *J. Am. Stat. Assoc.*, 63,
753 1379–1389, <https://doi.org/10.1080/01621459.1968.10480934>, 1968.
- 754 Sichangi, A. W., Wang, L., Yang, K., Chen, D., Wang, Z., Li, X., Zhou, J., Liu, W., and Kuria, D.:
755 Estimating continental river basin discharges using multiple remote sensing data sets, *Remote*
756 *Sens. Environ.*, 179, 36–53, <https://doi.org/10.1016/j.rse.2016.03.019>, 2016.
- 757 Smith, L. C.: Satellite remote sensing of river inundation area, stage, and discharge: A review,
758 *Hydrol. Process.*, 11, 1427–1439, [https://doi.org/10.1002/\(SICI\)1099-
759 \[1085\\(199708\\)11:10<1427::AID-HYP473>3.0.CO;2-S\]\(https://doi.org/10.1002/\(SICI\)1099-1085\(199708\)11:10<1427::AID-HYP473>3.0.CO;2-S\)](https://doi.org/10.1002/(SICI)1099-1085(199708)11:10<1427::AID-HYP473>3.0.CO;2-S), 1997.
- 760 Sun, M., Tangdamrongsub, N., Sun, Y., Dong, J., Sutanudjaja, E., and Smilovic, M.: Assessing
761 and optimizing high-resolution global river streamflow estimates with triple collocation
762 analysis, *J. Hydrol.*, 646, 135122, <https://doi.org/10.1016/j.jhydrol.2026.135122>, 2026.
- 763 Syvitski, J. P. M., Vörösmarty, C. J., Kettner, A. J., and Green, P.: Impact of humans on the flux
764 of terrestrial sediment to the global coastal ocean, *Science*, 308, 376–380,
765 <https://doi.org/10.1126/science.1109454>, 2005.
- 766 Tang, G., Clark, M. P., Knoben, W. J. M., Liu, H., Gharari, S., Arnal, L., Beck, H. E., Wood, A.
767 W., Newman, A. J., and Papalexiou, S. M.: The impact of meteorological forcing uncertainty
768 on hydrological modeling: A global analysis of cryosphere basins, *Water Resour. Res.*, 59,
769 e2022WR033767, <https://doi.org/10.1029/2022WR033767>, 2023.
- 770 Wang, C., Li, Z., Guyennon, N., Chen, Y., Li, Y., Liang, Q., and Di, Y.: Quantifying the
771 diminished glacier contribution to river runoff in the Tarim Basin, *Geophys. Res. Lett.*, 52,
772 e2025GL114641, <https://doi.org/10.1029/2025GL114641>, 2025.
- 773 Wang, K., Liu, X., Cui, P., Zhang, Y., Xie, J., Liu, C., and Gosling, S. N.: China's nationwide
774 streamflow decline driven by landscape changes and human interventions, *Sci. Adv.*, 11,
775 eadu8032, <https://doi.org/10.1126/sciadv.adu8032>, 2025a.



- 776 Wang, L., Li, X., Lutz, A., Nepal, S., Chen, D., Yao, T., Su, F., Cuo, L., Yao, Z., Zhang, Y., Hu,
777 Z., Huang, J., Hou, M., Liu, R., Long, J., Chai, C., Liu, Z., Bashir, A., Khanal, S., Sun, H.,
778 Nie, Y., Zhang, Y., and Wang, T.: Acceleration of diverging runoff trends on the Third Pole,
779 *Commun. Earth Environ.*, 6, 907, <https://doi.org/10.1038/s43247-025-02854-5>, 2025b.
- 780 Wang, Y. and Li, Y.: China Daily River Discharge Records (CDR²) dataset [Data set], Zenodo,
781 <https://doi.org/10.5281/zenodo.20152832>, 2026.
- 782 Xu, H.: Modification of normalised difference water index (NDWI) to enhance open water features
783 in remotely sensed imagery, *Int. J. Remote Sens.*, 27, 3025–3033,
784 <https://doi.org/10.1080/01431160600589179>, 2006.
- 785 Yang, L., Zhao, G., Tian, P., Mu, X., Tian, X., Feng, J., and Bai, Y.: Runoff changes in the major
786 river basins of China and their responses to potential driving forces, *J. Hydrol.*, 607, 127536,
787 <https://doi.org/10.1016/j.jhydrol.2022.127536>, 2022.
- 788 Yang, X., Pavelsky, T. M., Allen, G. H., and Donchyts, G.: RivWidthCloud: An automated Google
789 Earth Engine algorithm for river width extraction from remotely sensed imagery, *IEEE*
790 *Geosci. Remote Sens. Lett.*, 17, 217–221, <https://doi.org/10.1109/LGRS.2019.2920225>, 2019.
- 791 Yao, T., Bolch, T., Chen, D., Gao, J., Immerzeel, W., Piao, S., Su, F., Thompson, L., Wada, Y.,
792 Wang, L., Wang, T., Wu, G., Xu, B., Yang, W., Zhang, G., and Zhao, P.: The imbalance of
793 the Asian water tower, *Nat. Rev. Earth Environ.*, 3, 618–632, [https://doi.org/10.1038/s43017-](https://doi.org/10.1038/s43017-022-00299-4)
794 [022-00299-4](https://doi.org/10.1038/s43017-022-00299-4), 2022.
- 795 Yuan, Z., Lin, P., Guo, X., Zhang, K., and Beck, H. E.: Revisiting at-a-station hydraulic geometry
796 using discharge observations and satellite-derived river widths, *J. Remote Sens.*, 4, 0271,
797 <https://doi.org/10.34133/remotesensing.0271>, 2024.
- 798 Zheng, H., Yang, Z.-L., Lin, P., Wei, J., Wu, W.-Y., Li, L., Zhao, L., and Wang, S.: On the
799 sensitivity of the precipitation partitioning into evapotranspiration and runoff in land surface
800 parameterizations, *Water Resour. Res.*, 55, 95–111, <https://doi.org/10.1029/2017WR022236>,
801 2019.