



1 The Integrated Multi-proxy Paleoclimate Database for the Northeastern 2 Tibetan Plateau

3 Yu Li^{1,2}, Hao Shang¹, Shiyi Yan¹, Zhansen Zhang¹, MingJun Gao¹, Tingting Xi¹,
 4 Jiarui Zhao¹, Wanchen Sun¹

5 ¹College of Earth and Environmental Sciences, Lanzhou University, Lanzhou 730000,
 6 China

7 ²Key Laboratory of Western China's Environmental Systems (Ministry of Education),
 8 Lanzhou University, Lanzhou 730000, China

9 Correspondence: Yu Li (liyu@lzu.edu.cn)

10
 11 **Abstract** The Northeastern Tibetan Plateau (NETP) is an important region influenced
 12 by the interaction between the Asian monsoon and the Westerlies. Diverse
 13 paleoclimate archives, including sedimentary profiles, ice cores, tree rings, surface
 14 soils, and meteorological stations, record environmental changes across multiple
 15 temporal scales. However, differences in temporal resolution, chronological depth,
 16 and proxy sensitivity make direct comparisons among these archives difficult. We
 17 present a paleoclimate database for the NETP that integrates these records within a
 18 unified framework. The database includes 147 sedimentary profiles, 456 surface soil
 19 samples, 23 tree ring width index (RWI) chronologies, 5 ice core records, and data
 20 from 73 meteorological stations. The records were standardized using consistent
 21 chronological processing and proxy normalization methods to improve cross archive
 22 comparability. The database shows internal consistency among regional proxy records.
 23 The dataset combines the long term continuity of sedimentary profiles with the
 24 high-resolution climate signals preserved in tree rings and ice cores. In addition,
 25 integrating surface soil proxies with meteorological observations provides a modern
 26 calibration framework for evaluating proxy responses to temperature, precipitation,
 27 and elevation gradients. By linking geological archives with instrumental observations
 28 across multiple timescales, the dataset supports regional paleoclimate reconstruction
 29 and environmental change research in the NETP. The Integrated Multi-proxy
 30 Paleoclimate Database for the Northeastern Tibetan Plateau is available at
 31 <https://doi.org/10.17632/cp7kftnb94.1> (Li, 2026)

33 1 Introduction

34 Quantitative reconstruction of past climate change relies on a variety of
 35 paleoclimate proxies archived in diverse natural records, including ice cores, tree
 36 rings, sedimentary sequences, surface soils, and modern meteorological observations.
 37 These archives provide climatic information ranging from interannual to
 38 multi-millennial scales, enabling comprehensive analyses of climate variability across
 39 multiple temporal dimensions. With the continuous accumulation of observational
 40 data, the integrated application of multi-proxy indicators combining multiple proxies
 41 is now a key focus in paleoclimate research. However, proxies form and respond
 42 differently to the environment, resulting in varying temporal resolutions, climatic



meanings, and uncertainties. These differences limit dataset comparability and make integration difficult. This issue is especially prominent in regional studies involving complex environments.

As the highest plateau on Earth, the Tibetan Plateau (TP) is recognized as the "Third Pole" and an "amplifier of global climate change" (Pan et al., 1995; Wu et al., 2005; Qiu, 2008; Liu et al., 2009a). The Northeastern Tibetan Plateau (NETP) encompasses key geographical units, including the Qilian Mountains, the Qaidam Basin, the Hexi Corridor, and the Lanzhou section of the upper Yellow River. Notably, this region forms a critical intersection between the Asian monsoon and the Westerly jet stream. Such a unique climatic position provides distinct advantages for recording paleoclimatic evolution (Long et al., 2025; Li et al., 2025; Wang et al., 2026). Extensive research has established the NETP as one of the most mature regions globally for paleoclimate studies since the Last Glacial Maximum (LGM). It hosts the richest variety of proxy types in the global mid to high latitudes, including surface soils (Wang et al., 2017; Chen et al., 2020; Zan et al., 2022; Li and Peng, 2023), sedimentary profiles (Gasse et al., 1991; Liu et al., 2014; Wünnemann et al., 2010; Herzschuh et al., 2011; Kasper et al., 2012; Bird et al., 2014; Doberschütz et al., 2014; Hou et al., 2017a, 2017b; Conroy et al., 2017;), tree rings (Yang et al., 2012; Gou et al., 2013; He et al., 2013; Qie et al., 2022), and ice cores (Thompson et al., 1997; Yao et al., 1997). This dense network helps track spatial and temporal changes of climate elements.

Despite this data wealth, integrating these diverse carriers remains challenging. Meteorological stations provide high-precision modern records; tree rings offer clear interannual resolution; ice cores record long-sequence climatic information; and stratigraphic profiles provide continuous paleoenvironmental sequences over multi-millennial scales. These differences in spatio-temporal resolution make direct comparisons difficult. Also, differing proxy responses add to this complexity. For instance, early Holocene lacustrine carbonate $\delta^{18}\text{O}$ records from Lake Qinghai are typically interpreted as indicating increased effective moisture and a strengthened summer monsoon (Lister et al., 1991; Liu et al., 2007; An et al., 2012). However, lake-level proxies and paleoshoreline evidence from the same lake indicate a relatively dry climate during this period (Zhang et al., 1994; Liu et al., 2013; Wang et al., 2014; Liu et al., 2015). Holocene temperature variations also exhibit discrepancies between pollen-inferred early to mid Holocene warming (Tang et al., 1999, 2000, 2009; Herzschuh et al., 2006; Shen et al., 2008; Lu et al., 2011; Ma et al., 2014) and the mid Holocene cooling phase recorded by the Guliya ice core $\delta^{18}\text{O}$ (Thompson et al., 1997). To address these limitations, this study constructs a regional multi-proxy paleoclimate database for the NETP. By systematically integrating sedimentary profiles, surface soils, tree rings, ice cores, and meteorological observations, the database integrating these records unifies the diverse data sources. All sedimentary records were converted into a unified chronological sequence using age-depth models and interpolated to a consistent temporal resolution. Subsequently, Z-score normalization was applied to proxy indicators to eliminate dimensional effects. This framework allows for cross-proxy comparisons and quantitative analysis, and

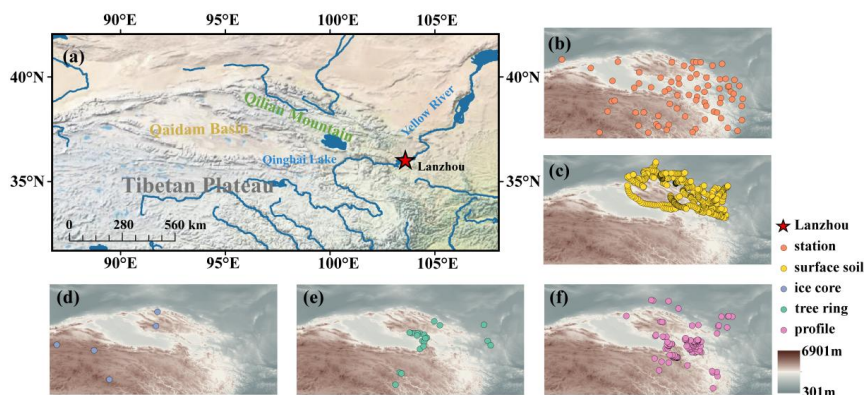


87 provides constraints for climate model validation.

88

89 2 Study region

90 The study region is located on the Northeastern Tibetan Plateau (98.0°-103.5°E,
91 32.7°-38.2°N), situated in the transition zone between China's first and second
92 topographic steps, encompassing typical geographical units such as the Qaidam Basin,
93 the Hexi Corridor, the Lanzhou section of the upper Yellow River, and the Qilian
94 Mountains (Fig. 1). Regional elevation generally ranges between 2 and 4 km and
95 gradually rises from the basins and corridor areas to over 5000 m in the mountains,
96 forming a significant vertical altitudinal gradient. The complex terrain causes high
97 spatial variation in sedimentary and hydrological systems. It also creates gaps in the
98 climate records in the paleoclimatic information recorded by various carriers. The
99 NETP is jointly influenced by the East Asian Summer Monsoon (EASM), the Indian
100 Summer Monsoon (ISM), and the Westerly circulation, making it extremely sensitive
101 to global changes (Ding and Wang, 2005; Schiemann et al., 2009; Chen et al., 2020;
102 Huang et al., 2020; Li et al., 2022). Since the Late Glacial, the interplay and
103 fluctuation of these different circulation systems have resulted in strong
104 spatio-temporal asymmetry in regional hydrothermal configurations. This increases
105 the uncertainty of single-proxy reconstructions and complicates data integration
106 across different terrains.



107

108 Figure 1. Spatial distribution of multi-proxy climate records in the Northeastern
109 Tibetan Plateau. (a) Extent and geographical units of the Northeastern Tibetan
110 Plateau (NETP); (b-f) spatial distribution of meteorological stations, surface soils,
111 tree rings, and sedimentary profiles integrated in this database.

112

113 3 Data compilation and processing

114 3.1 Data compilation

115 This study integrates five categories of data: sedimentary profiles, surface soil
116 samples, tree ring records, ice core records, and meteorological stations (Tab. 1). This
117 database covers timescales from modern observations to millennia and includes
118 various physical and geochemical proxies. Sedimentary profile data were acquired
119 through systematic literature searches across databases including Web of Science,



Elsevier, CNKI, and Google Scholar, with studies containing original chronological information specifically selected. A total of 147 profiles were integrated, including metadata such as location, elevation, depth, and dating materials. This yielded a compilation of 1,006 chronological data points and multi-proxy sequences, including grain size, total organic carbon (TOC), magnetic susceptibility, C/N, $\delta^{13}\text{C}$, $\delta^{18}\text{O}$, as well as carbonates, silicates, and Rb/Sr ratios. Surface soil data were derived from field sampling and laboratory analysis, comprising 456 samples with location and elevation information, providing records of grain size, C/N ratios, and 289 carbon and oxygen stable isotope values. tree ring data were obtained from the NOAA database, consisting of 23 tree ring width index (RWI) chronologies, along with information on tree species and their distributions. Ice core data integrate five records from Dunde, Malan, Tanggula, Muztagh, and Laohugou, sourced from the NOAA, China Tibetan Plateau Data Center (TPDC) and published literature, providing $\delta^{18}\text{O}$ sequences and selected ionic proxies. Meteorological observation data were sourced from NOAA, containing temperature and precipitation records from 73 stations for the period 1973-2025.

Proxy type	Number of records	Temporal coverage (CE)	Spatial coverage	Elevation range (m)	Parameters
Station	73	1973-2025	32.20-39.00°N、 85.60-105.00°E	609-4877	MAT, MAP
Surface soil	456	-	36.17-41.35°N、 94.3-103.64°E	1037-415 8	grain size, $\delta^{13}\text{C}_{\text{org}}$, $\delta^{18}\text{O}$, C/N, temperature , precipitation
Ice core	5	520-2010	33.11-39.43°N、 87.179-96.57°E	5049-575 0	$\delta^{18}\text{O}$, $\text{C}_2\text{O}_4^{2-}$, SO_4^{2-} , Ca^{2+}
Tree ring	23	-336-2020	32.66-38.57°N、 95.72-108.90°E	1050-432 4	RWI
Profile	147	-28050-2025	32.33-41.89°N、 94.00-104.00°E	892-5133	grain size, TOC, magnetic susceptibility, C/N, $\delta^{13}\text{C}$, $\delta^{18}\text{O}$, OM, CaCO_3 , carbonates, silicates , Rb/Sr

Table 1. Dataset Overview

3.2 Data processing

3.2.1 Processing of sedimentary records

This study compiled 1,006 published dating records to establish temporal frameworks for the sedimentary profiles. Radiocarbon dates were calibrated to calendar ages using the IntCal20 calibration curve. The median probability age was used as the chronological control point. For optically stimulated luminescence (OSL) records, the original text data were cleaned to extract representative age values. To reduce the influence of possible depositional disturbance, age inversions were



145 identified and excluded before further processing. All original dating information and
146 uncertainties were preserved in the released database. The retained dating points were
147 used to establish approximate age-depth relationships for each profile. Simple fitted
148 relationships, including linear and low-order polynomial functions, were applied to
149 support temporal interpolation and alignment among different records. Because the
150 compiled profiles contain different dating methods and uneven chronological
151 resolution, the fitted relationships were intended to provide generalized temporal
152 support rather than high-precision chronologies. The fitted results do not replace the
153 original dating data, which remain fully available in the released database. To support
154 comparisons among different profiles, all proxy records were converted from the
155 depth domain to the time domain. Each sequence was resampled at 50-year intervals
156 to achieve a uniform temporal resolution. The corresponding depths for target ages
157 were calculated using the fitted age-depth relationships. Proxy values were then
158 obtained through linear interpolation. This procedure generated evenly spaced time
159 series for subsequent regional analyses. A total of 45 profiles with relatively complete
160 records and continuous proxy data were selected for integrated analysis. Each proxy
161 sequence was standardized independently using Z-score normalization based on its
162 own mean and standard deviation (Fig. 2). Ensemble statistics were then calculated
163 for the standardized records at the same time intervals, including regional means and
164 standard deviations. These results were used to describe overall environmental
165 variability and data dispersion across the study region.

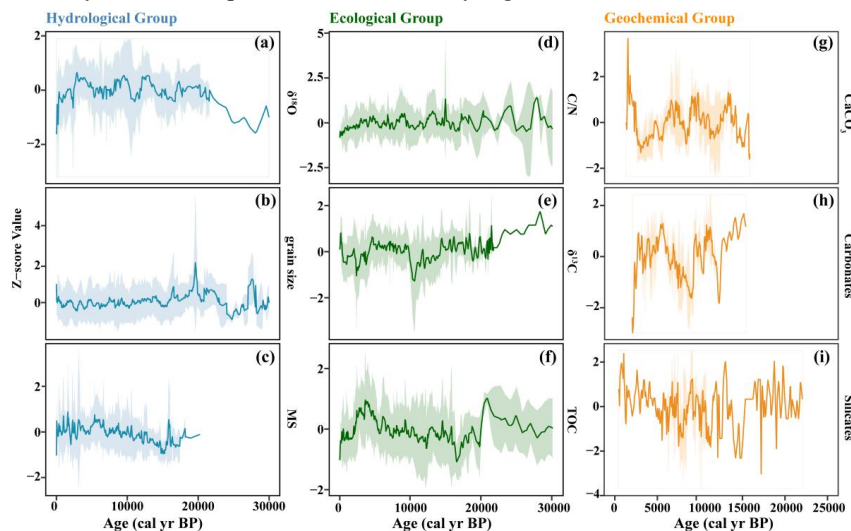


Figure 2. Standardized proxy time series from sedimentary records

3.2.2 Processing of surface soil dataset

Surface soil data in this study were collected from 456 sampling sites across various altitudinal and environmental gradients to ensure regional representativeness. All samples were processed and analyzed at the Key Laboratory of Western China's Environmental Systems (Ministry of Education), Lanzhou University. Grain size analysis was performed using a Malvern Mastersizer 2000 laser diffraction particle



size analyzer, with a measurement range of 0.02-2000 μm and a repetitive error restricted to within 3%. The Carbon to Nitrogen (C/N) ratio was calculated by measuring organic carbon (OC) and total nitrogen (TN) contents via an elemental analyzer. Stable carbon and oxygen isotopic compositions were determined using an isotope ratio mass spectrometer (IRMS), with $\delta^{13}\text{C}_{\text{org}}$ and $\delta^{18}\text{O}$ values calibrated against the Vienna Pee Dee Belemnite (VPDB) and Vienna Standard Mean Ocean Water (VSMOW) standards, respectively. Additionally, the corresponding Mean Annual Temperature (MAT) and Mean Annual Precipitation (MAP) for each sampling site were obtained through spatial matching.

3.2.3 Processing of ice core records

This study integrated five ice core records from Dundee, Malan, Tanggula, Muztagh, and Laohugou, sourced from the National Oceanic and Atmospheric Administration (NOAA) and published literature. Significant variations exist in the temporal coverage and resolution among these records; the Dundee ice core provides the longest temporal span but with a decadal raw resolution, whereas the Tanggula and Muztagh cores consist of annual-resolution data. Records for the Malan and Laohugou cores were derived from published studies. For sequences sourced from literature with inconsistent resolutions, linear interpolation was applied to unify the data to an annual scale. The processed ice core data preserve their original variation patterns under a unified temporal scale.

3.2.4 Processing of tree ring and instrumental data

Tree ring data were retrieved from the NOAA public database, consisting of tree ring width index (RWI) chronologies within the study area at an annual resolution. The RWI sequences used in this study were the standardized products provided by the database, and no additional processing was performed. Meteorological observation data were likewise obtained from NOAA, encompassing daily temperature and precipitation records from 73 stations in the study area from 1973 to 2025. The original units were converted from Fahrenheit and inches to Celsius and millimeters, respectively, to ensure data consistency and comparability. Daily observations were aggregated into annual sequences.

4 Results and discussions

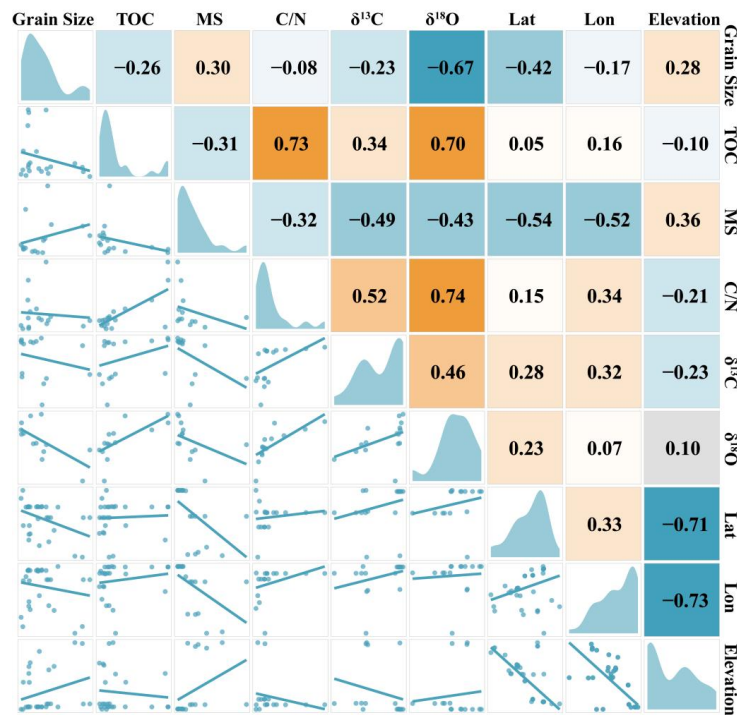
4.1 Proxy relationships

Sedimentary profiles provide continuous, multi-proxy records for reconstructing paleoclimates. (Chen et al., 2020). Organic geochemical proxies and micro-biomarkers in lacustrine sediments can reconstruct the evolution of regional ecosystem productivity, large oscillations in summer temperatures, and the temperature-driven drying processes in inland basins (Hou et al., 2016; Wu et al., 2020). Physical characteristics in aeolian and loess strata are vital indicators for tracking dust transport dynamics, the intensity evolution of the East Asian winter monsoon, and the history of regional aeolian desertification (Lehmkuhl et al., 2014; Li and Morrill, 2015; Hu et al., 2017, 2018).

Correlation analysis (Fig. 3) based on profile data with multi-proxy records demonstrates the statistical relationships among different proxy indicators and their



218 associations with geographical factors. Correlation strengths vary, showing
219 inconsistent links between proxies. Among geographical factors, elevation exhibits
220 relatively strong negative correlations with both latitude and longitude, with the
221 corresponding scatter distributions showing an overall downward trend; however, the
222 point distributions still possess a certain degree of dispersion and do not present a
223 strictly linear relationship. $\delta^{18}\text{O}$ exhibits a significant negative correlation with grain
224 size. Furthermore, $\delta^{18}\text{O}$ shows significant positive correlations with TOC and C/N,
225 characterized by relatively concentrated scatter distributions. Magnetic susceptibility
226 (MS) presents relatively significant negative correlations with $\delta^{13}\text{C}$, $\delta^{18}\text{O}$, longitude,
227 and latitude, with its scatter distributions being relatively dispersed. In contrast, $\delta^{18}\text{O}$,
228 TOC, and C/N exhibit weak correlations with environmental factors, including
229 elevation, longitude, and latitude.



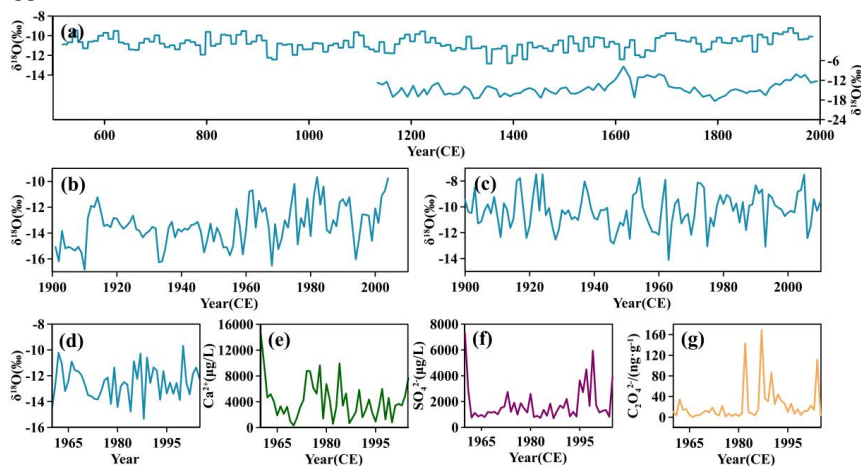
230
231 Figure 3. Correlation among proxy variables and geographic factors. (Note: MS
232 denotes magnetic susceptibility.)

233 4.2 Temporal characteristics

234 Ice cores provide strong sedimentary continuity, high dating precision, and the
235 coexistence of multiple proxy indicators, constitute crucial data for multi-timescale
236 environmental reconstructions on the Tibetan Plateau (Cheng, 2026). $\delta^{18}\text{O}$ tracks the
237 regional water cycle and helps delineate the boundary of influence between the Asian
238 summer monsoon and the Westerlies (Tian et al., 2001; Wang et al., 2006). Pollen
239 assemblages and specific chemical microparticles preserved in ice cores show past
240 regional vegetation changes and surface material emissions (Liu et al., 1998; Cui et al.,



241 2011a; Cui et al., 2011b). As shown in Fig. 4, the five ice core sequences (Dunde,
 242 Malan, Muztagh, Tanggula, and Laohugou) integrated in this dataset can be classified
 243 based on their temporal coverage. Dunde and Malan represent long timescale records;
 244 Muztagh and Tanggula provide centennial scale observations; while Laohugou serves
 245 as a high-resolution, 46-year short sequence. Across all records, Large $\delta^{18}\text{O}$ variations
 246 across all records indicate strong climate fluctuations. The long-timescale cores
 247 distinctly capture historical cold phases likely corresponding to the Little Ice Age
 248 (LIA). The Dunde ice core exhibits pronounced $\delta^{18}\text{O}$ depletion during the 14th and
 249 17th centuries. The Malan sequence records a more severe cooling episode, with
 250 values plunging to an absolute minimum of -18.3 in the 1790s. The Malan record also
 251 displays the widest isotopic range (-18.3 to -7.9), reflecting extreme climate
 252 fluctuations. In the 20th century, the records show different trends. The Malan
 253 sequence exhibits a continuous, long-term warming trajectory from 1900 to 1994. The
 254 Dunde record peaked in the 1940s before experiencing a reversal in the 1960s and
 255 1970s. This mid to late 20th-century cooling phase is concurrently recorded as distinct
 256 low-value troughs in both the Muztagh and Tanggula centennial sequences between
 257 the 1950s and 1970s. In recent decades, multiple cores indicate an intensified
 258 warming signal, marked by positive $\delta^{18}\text{O}$ anomalies in the Muztagh and Tanggula
 259 records during the early 2000s. The short Laohugou sequence captures
 260 extreme environmental events. Its $\text{C}_2\text{O}_4^{2-}$ concentrations generally remained at low
 261 background levels, but showed sudden spikes in 1982, 1987, and 2004. These peaks
 262 suggest sudden environmental events.



263 Figure 4. Multi-timescale ice core proxy records. (a) Long timescale Dunde and
 264 Malan ice cores; (b-c) medium timescale Tanggula and Muztagh ice cores; (d-g) short
 265 timescale Laohugou ice core.

267 Tree rings provide precise interannual resolution and are highly sensitive to
 268 environmental changes. This makes them key proxies for reconstructing the climate of
 269 the Northeastern Tibetan Plateau over the past millennium. (Hou et al., 1999).
 270 However, the response relationship between radial tree growth and climatic elements
 271 is not singular. It is driven by interactions between climate and species-specific traits.



272 Even within the same type of habitat, the climate signals captured by different tree
 273 species exhibit significant differences: for instance, in the eastern Qilian Mountains,
 274 the ring width of Qilian juniper primarily responds to temperature fluctuations,
 275 whereas Qinghai spruce is more sensitive to precipitation changes (Zhan et al., 2019).
 276 Furthermore, the climate response patterns of the same tree species diverge across
 277 different altitudinal and latitudinal gradients (Liang et al., 2010; Yang et al., 2013).
 278 This sensitivity rule has allowed tree ring width to be widely and successfully applied
 279 in reconstructing regional large-scale precipitation fluctuations (Gou et al., 2013; He
 280 et al., 2013), historical runoff inversion (Liu et al., 2009b; Yang et al., 2012), and the
 281 precise identification of extreme flood events (Qie et al., 2022). This dataset includes
 282 23 high quality RWI chronologies to capture these complex responses (Fig. 5). High
 283 altitude Qilian juniper sequences provide long and stable climate signals. Low altitude
 284 Chinese pine sequences are highly sensitive to abrupt local environmental changes.
 285 Under a unified annual resolution framework, Qilian juniper sampling sites
 286 distributed in high-altitude regions generally possess ultra-long temporal coverages
 287 ranging from hundreds to thousands of years, which is attributed to the extreme
 288 natural longevity of this species (Yang et al., 2014; Zhang et al., 2014; Yin et al.,
 289 2016). High altitude Qilianshan juniper chronologies exhibit relatively stable radial
 290 growth trajectories. Their RWI series demonstrate low-amplitude variance and
 291 attenuated low-frequency fluctuations. Conversely, low altitude Chinese pine
 292 sequences are characterized by pronounced interannual volatility, with their RWI
 293 series displaying extreme positive and negative anomalies. Statistical evaluation
 294 reveals that the variance amplitude of the pine chronologies is approximately three
 295 times that of the juniper records. This indicates that low altitude forest ecosystems are
 296 sensitive to short term, localized hydroclimatic extremes. In contrast, the dampened
 297 variance of the high altitude juniper sequences more reliably captures long term, large
 298 scale regional climate forcing.

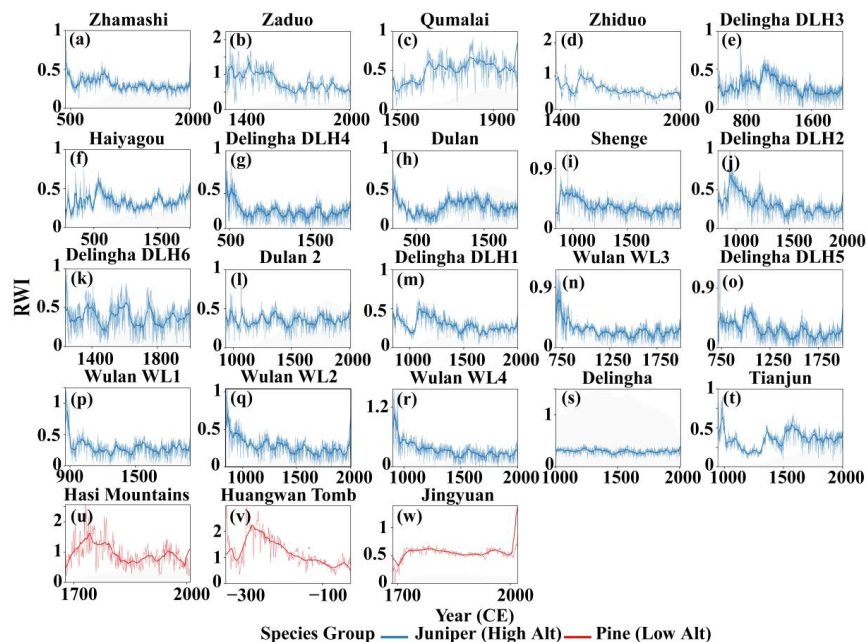


Figure 5. tree ring width index (RWI) chronologies across elevation gradients. (a-t) RWI chronologies of Qilian juniper (*Sabina przewalskii* Kom.) from high-altitude regions in the Northeastern Tibetan Plateau (NETP); (u-w) RWI chronologies of Chinese pine (*Pinus tabuliformis* Carrière) from low-altitude regions in the NETP.

4.3 Modern observations and proxy context

Station records and surface soils form the modern baseline for this database. (Shi et al., 2022). Comparing surface soil proxies with station data reveals how sedimentary proxies respond to regional climate gradients. (Li and Peng, 2023; Wang et al., 2017). In the monsoon marginal zone, the significant positive correlation between surface soil magnetic characteristics and station-observed precipitation helps delineate the boundaries between semi arid and arid climates (Zan et al., 2022). The spatial heterogeneity of surface soil grain-size distributions clearly records the interaction between modern weathering dynamics and local hydrological processes (Chen et al., 2020). However, in extremely arid inland interiors, some soil proxies do not match precipitation records, suggesting that non climatic factors must be fully considered when interpreting proxy records in specific environments (Huo et al., 2025). This calibration validates the proxies and provides a modern baseline to isolate non-climatic noise for calibrating high-resolution records. (Li and Peng, 2023; Wang et al., 2017). Surface soil data (Fig. 6) quantitatively characterize the statistical relationship structure between proxy indicators and modern climatic variables and topographic factors. From the regression results, the C/N ratio exhibits a significant negative correlation with precipitation, a weak positive correlation with temperature, and a weak correlation with elevation, indicating variations in its response intensity to different climatic variables. Grain size displays consistent trends across three dimensions: decreasing with increasing elevation, decreasing with increasing



precipitation, and increasing with increasing temperature. $\delta^{13}\text{C}$ is negatively correlated with precipitation, while exhibiting weak relationships with temperature and elevation; the correlation of $\delta^{18}\text{O}$ across all variable directions is generally low. Regarding distribution characteristics, the sampling sites are primarily concentrated in two high-density intervals—temperatures of 5–10 °C, precipitation of 50–150 mm, and elevations of 1000–1500 m and 3000–3500 m—revealing a distinct spatial stratification structure. In the scatter distributions, most data points cluster around the regression trends, with a limited proportion of outliers, indicating the overall stability of the statistical relationships.

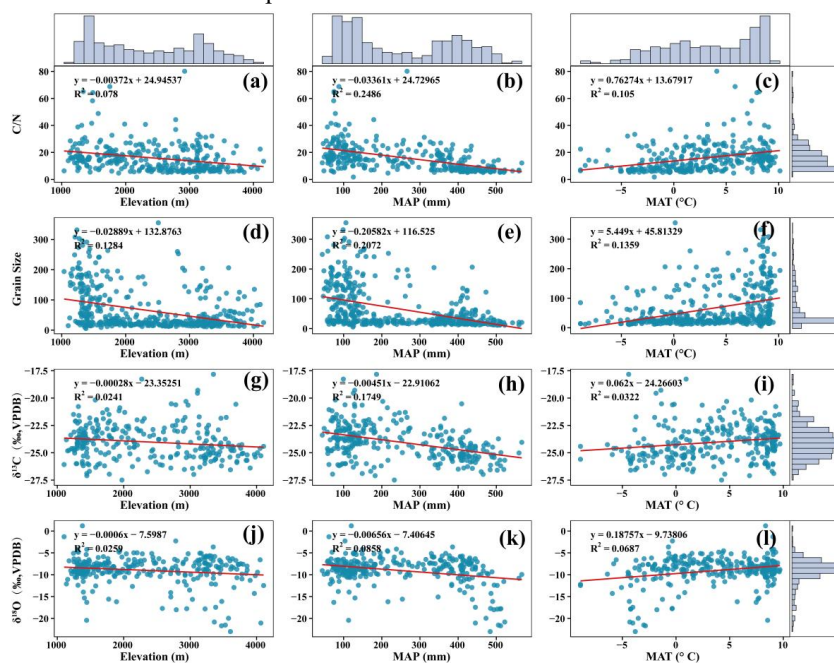


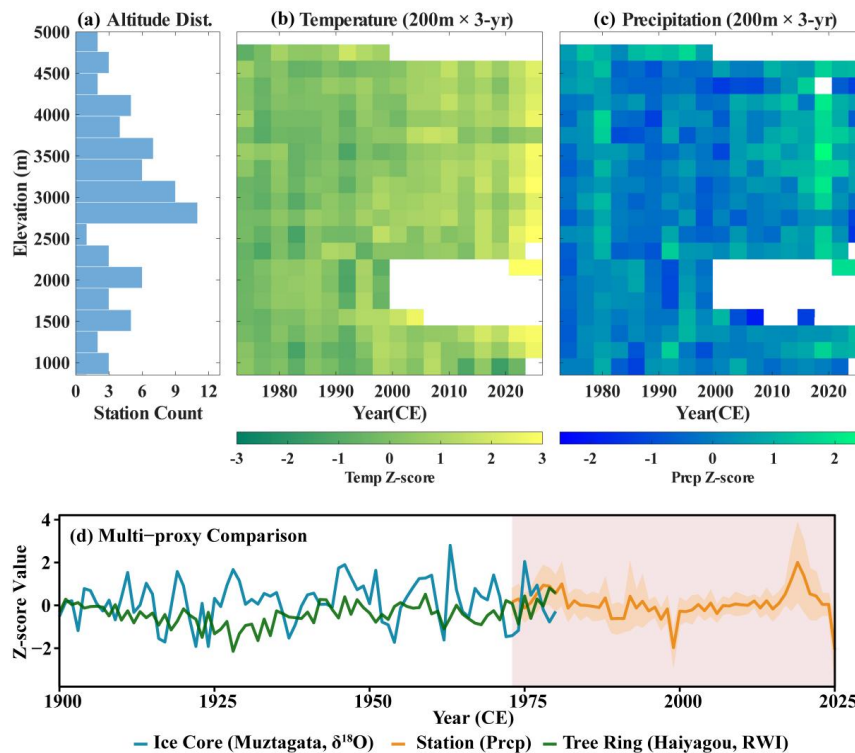
Figure 6 Surface soil proxy-environment relationships

Meteorological station data provide modern climate baselines (Fig. 7). The stations are primarily distributed within the altitudinal range of 2600–3500 m, with spatial coverage concentrated in the mid to high altitude zones. Z-score heatmaps based on 200 m elevation bins and 3-year averages indicate that temperature was dominated by negative anomalies during most periods, with the extent of positive anomalies expanding significantly after 2015; during the same period, precipitation also exhibited episodic characteristics of transitioning from negative to positive anomalies. Different altitudinal zones displayed similar anomaly signs during the same periods, although their intensities varied.

In the multi-archive comparison (Fig. 7_d), the Muztagata ice core $\delta^{18}\text{O}$ and tree ring width indices (RWI) provide medium to long timescale information (1900–1980), while the precipitation data recorded by stations cover the modern stage (1973–2025). Since high original ice core $\delta^{18}\text{O}$ values indicate arid environments, which physically contrasts with the humid conditions indicated by high RWI and instrumental



precipitation, the $\delta^{18}\text{O}$ sequence was uniformly multiplied by -1 to facilitate direct comparison. After unifying the environmental indications, the overall variation trends of the $\delta^{18}\text{O}$ and RWI between 1900 and 1980 are similar, showing a downward trend during 1913-1919 and 1961-1967, and an upward trend during 1967-1969. However, these two paleoclimate sequences demonstrated pronounced opposite trends during the 1920s and 1930s. Furthermore, a comparison of their fluctuation characteristics reveals that the overall variations of the RWI generally lag the $\delta^{18}\text{O}$ sequence by approximately 4 to 5 years. During the overlapping period with station precipitation data (1973-1980), the RWI, $\delta^{18}\text{O}$, and station precipitation data maintained consistency in their overall variation trends. This trend consistency during the overlapping period establishes an effective temporal calibration, logically allowing the variations of paleoclimate proxies to extend into modern instrumental records. Beyond the splicing of time series, the sedimentary profile and topsoil (Fig. 3) further perfect the continuous record of regional environmental evolution. On this basis, the co-variation of multiple profile proxies within a unified temporal framework reflects the integrated response of this record to multiple environmental factors.



366

367

368 Figure 7. Modern observations and cross-archive continuity. (a) Histogram of station
 369 distribution across elevation gradients; (b-c) Heatmaps showing the spatiotemporal
 370 variations of station-observed temperature and precipitation across elevation and time;
 371 (d) Comparison of multi-archive proxies including ice core, tree ring, and
 372 meteorological stations, where the orange line represents the standardized (Z-score)



multi-year precipitation observations from stations.

4.4 Challenges and future directions of paleoclimate research in the NETP

Integrating diverse records in the NETP is challenging due to inherent differences in temporal resolution. Meteorological station data and tree rings provide precise annual or sub-annual records. In contrast, ice cores and sedimentary profiles generally reflect decadal to centennial variations. Furthermore, surface soil samples primarily offer modern spatial baselines rather than continuous time series. Disparities in chronological control methods across different archives further complicate integration. Sedimentary profiles rely on radiocarbon (AMS ^{14}C) and optically stimulated luminescence (OSL) dating to establish age-depth models, tree ring sequences provide absolute chronological constraints through cross-dating, and ice core records are calibrated based on layer counting or numerical modeling. The inherent differences in precision and uncertainty estimation among these methods may lead to temporal offsets or inconsistencies across multiple sequences (Reimer et al., 2020; Bronk Ramsey, 2009). At the proxy level, parameters such as $\delta^{18}\text{O}$, grain size, and C/N ratios respond to hydrological processes and associated climatic conditions; however, their numerical ranges and environmental sensitivities vary, which may introduce additional uncertainties during the unified standardization process, particularly in cross regional or altitudinal comparisons (Bradley, 2015). Spatially, constrained by accessibility and preservation conditions in high-altitude terrain, existing data tend to cluster in more accessible river valleys or piedmont areas, resulting in a relative paucity of sampling sites in high-altitude transition zones and key watersheds. This limits our ability to map temperature and precipitation gradients. Regarding modern observations, although meteorological station records provide high temporal resolution, their coverage is primarily limited to recent decades, exhibiting only narrow overlap with tree ring and ice core records. This restricts comparisons to short time windows and limits the direct validation of long-term variations (PAGES 2k Consortium, 2017).

This multi-archive dataset integrates records by standardizing diverse proxies into a single framework. Future research necessitates higher-resolution sampling in key regions to reduce discrepancies in temporal intervals, alongside improvements in age-modeling and uncertainty quantification to enhance temporal alignment precision across sequences. Furthermore, with the advancement of data assimilation and machine learning, coupling multi-source proxy data with climate models has become a pivotal direction for exploring proxy relationships (Hakim et al., 2016; Dee et al., 2016). The continuous optimization of sampling resolution, chronological frameworks, and proxy systems, as well as the systematic integration of multi-archive data within a unified framework can enhance the development of paleoclimate research in this region.

5 Data availability

The Integrated Multi-proxy Paleoclimate Database for the Northeastern Tibetan Plateau is now publicly available, and the data can be found at <https://doi.org/10.17632/cp7kftnb94.1> (Li, 2026)



417

418 6 Conclusion

419 This study constructs the Integrated Multi-proxy Paleoclimate Database for the
 420 Northeastern Tibetan Plateau integrating sedimentary profiles, surface soils, tree rings,
 421 ice cores, and meteorological stations under a unified chronological and Z-score
 422 standardized framework. Integration of multiple archives differentiates climate signals
 423 across timescales. Sedimentary profiles provide continuous sequences on a millennial
 424 scale. Deep ice cores and high altitude tree rings track low frequency background
 425 shifts. Short ice core records and low altitude tree rings capture high frequency
 426 environmental extremes. Cross validation between surface soils and modern
 427 meteorological stations establishes a spatial calibration baseline. This clarifies the
 428 sensitivity of proxy indicators to temperature, precipitation, and elevation gradients.
 429 By bridging geological archives with instrumental observations, this database
 430 provides fundamental data for regional multiple proxy comparisons and multiple
 431 timescale paleoclimate reconstructions. This integrated dataset provides a long term
 432 benchmark for understanding environmental evolution in this sensitive global climate
 433 change amplifier. Furthermore, these standardized records offer critical observational
 434 constraints for validating paleoclimate models and predicting regional climate
 435 dynamics.

436

437 Author contributions. YL: conceptualization and investigation. HS: data collection,
 438 data processing and writing (original draft preparation). SY: methodology and writing
 439 (reviewing and editing). ZZ and MG: software and data curation. TX, JZ, and WS:
 440 validation and resources.

441

442 Competing interests. The contact author has declared that none of the authors has any
 443 competing interests.

444

445 Acknowledgements.

446 We wish to thank the editor and the anonymous reviewers for providing critical and
 447 valuable comments that helped to improve the paper. We thank all of the researchers
 448 and institutions that provided data. During the preparation of this manuscript, we used
 449 Gemini 3 to improve readability. After using this service, we carefully and thoroughly
 450 reviewed and revised the content as necessary. We take full responsibility for the
 451 integrity and accuracy of the publication.

452

453 Financial support. This research was supported by the National Natural Science
 454 Foundation of China (42371159).

455

456 References

457 An, Z., Colman, S. M., Zhou, W., Li, X., Brown, E. T., Jull, A. J. T., Cai, Y., Huang, Y.,
 458 Lu, X., Chang, H., Song, Y., Sun, Y., Xu, H., Liu, W., Jin, Z., Liu, X., Cheng, P.,
 459 Liu, Y., Ai, L., Li, X., Liu, X., Yan, L., Shi, Z., Wang, X., Wu, F., Qiang, X.,
 460 Dong, J., Lu, F., and Xu, X.: Interplay between the westerlies and asian monsoon



- 461 recorded in lake Qinghai sediments since 32 ka, *Sci. Rep.*, 2, 619,
 462 <https://doi.org/10.1038/srep00619>, 2012.
- 463 Bird, B. W., Polisar, P. J., Lei, Y., Thompson, L. G., Yao, T., Finney, B. P., Bain, D. J.,
 464 Pompeani, D. P., and Steinman, B. A.: A Tibetan lake sediment record of
 465 Holocene Indian summer monsoon variability, *Earth Planet. Sci. Lett.*, 399,
 466 92-102, <https://doi.org/10.1016/j.epsl.2016.09.004>, 2014.
- 467 Bradley, R. S.: *Paleoclimatology: Reconstructing Climates of the Quaternary*, 3rd edn.,
 468 Academic Press, 2015.
- 469 Bronk Ramsey, C.: Bayesian analysis of radiocarbon dates, *Radiocarbon*, 51, 337-360,
 470 <https://doi.org/10.1017/S0033822200033865>, 2009.
- 471 Chen, F., Zhang, J., Liu, J., Cao, X., Hou, J., Zhu, L., Xu, X., Liu, X., Wang, M., Wu,
 472 D., Huang, L., Zeng, T., Zhang, S., Huang, W., Zhang, X., and Yang, K.: Climate
 473 change, vegetation history, and landscape responses on the Tibetan Plateau
 474 during the Holocene: A comprehensive review, *Quat. Sci. Rev.*, 243, 106444,
 475 <https://doi.org/10.1016/j.quascirev.2020.106444>, 2020.
- 476 Cheng, H.: The age of ice cores from Qinghai-Xizang Plateau, *Chin. Sci. Bull.*,
 477 <https://doi.org/10.1360/CSB-2025-5593>, 2026.
- 478 Conroy, J. L., Hudson, A. M., Overpeck, J. T., Liu, K., Wang, L., and Cole, J. E.: The
 479 primacy of multidecadal to centennial variability over late-Holocene forced
 480 change of the Asian Monsoon on the southern Tibetan Plateau, *Earth Planet. Sci.*
 481 *Lett.*, 458, 337-348, <https://doi.org/10.1016/j.epsl.2016.10.044>, 2017.
- 482 Cui, X., Ren, J., Qin, X., Du, W., Sun, W., Yu, G., Wang, Z., Wang, X., Sun, W., and
 483 Wang, Y.: Climatic and Environmental Records within a Shallow Ice Core at
 484 Laohugou Glacier No.12, Qilian Mountains, *J. Glaciol. Geocryol.*, 33,
 485 1251-1258, <https://doi.org/10.7522/j.issn.1000-0240.2011.0171>, 2011a (in
 486 Chinese).
- 487 Cui, X., Ren, J., Qin, X., Yu, G., Liu, W., and Xie, A.: Oxalate, fluoride record and
 488 their environmental significance in Laohugou Glacier 12 Qilian Mountains,
 489 *Environ. Chem.*, 30, 1919-1925,
 490 [https://www.sciengine.com/EC/articleIndex?doi=10.0000/j.0254-6108.20113011](https://www.sciengine.com/EC/articleIndex?doi=10.0000/j.0254-6108.2011301119191925&scroll=)
 491 [19191925&scroll=](https://www.sciengine.com/EC/articleIndex?doi=10.0000/j.0254-6108.2011301119191925&scroll=), 2011b (in Chinese).
- 492 Dee, S. G., Steiger, N. J., Emile-Geay, J., and Hakim, G. J.: On the utility of proxy
 493 system models for estimating climate states over the Common Era, *J. Adv.*
 494 *Model. Earth Syst.*, 8, 1164-1179, <https://doi.org/10.1002/2016MS000677>, 2016.
- 495 Ding, Q. and Wang, B.: Circumglobal teleconnection in the northern hemisphere
 496 summer, *J. Climate*, 18, 3483-3505, <https://doi.org/10.1175/jcli3473.1>, 2005.
- 497 Doberschütz, S., Frenzel, P., Haberzettl, T., Kasper, T., Wang, J., Zhu, L., Daut, G.,
 498 Schwalb, A., and Mausbacher, R.: Monsoonal forcing of Holocene
 499 paleoenvironmental change on the central Tibetan Plateau inferred using a
 500 sediment record from Lake Nam Co (Xizang, China), *J. Paleolimnol.*, 51,
 501 253-266, <https://doi.org/10.1007/s10933-013-9702-1>, 2014.
- 502 Gasse, F., Arnold, M., Fontes, J. C., Fort, M., Gibert, E., Huc, A., Li, B., Li, Y., Liu,
 503 Q., and Melieres, F.: A 13,000-year climate record from western Tibet, *Nature*,
 504 353, 742-745, <https://doi.org/10.1038/353742a0>, 1991.



- 505 Gou, X., Yang, T., Gao, L., Deng, Y., Yang, M., and Chen, F.: A 457-year
 506 reconstruction of precipitation in the southeastern Qinghai Tibet Plateau, China
 507 using tree-ring records, *Chin. Sci. Bull.*, 58, 978-985,
 508 <https://doi.org/10.1007/s11434-012-5539-7>, 2013 (in Chinese).
- 509 Hakim, G. J., Emile-Geay, J., Steig, E. J., et al.: The Last Millennium Climate
 510 Reanalysis project: Framework and first results, *J. Geophys. Res.-Atmos.*, 121,
 511 6745-6764, <https://doi.org/10.1002/2016JD024751>, 2016.
- 512 He, M., Yang, B., Bräuning, A., Wang, J., and Wang, Z.: Tree-ring derived millennial
 513 precipitation record for the south-central Tibetan Plateau and its possible driving
 514 mechanism, *Holocene*, 23, 36-45, <https://doi.org/10.1177/0959683612450198>,
 515 2013.
- 516 Herzschuh, U.: Palaeo-moisture evolution in monsoonal Central Asia during the last
 517 50,000 years, <https://doi.org/10.1016/j.quascirev.2005.02.006>, *Quat. Sci. Rev.*, 25,
 518 163-178, 2006.
- 519 Herzschuh, U., Ni, J., Birks, H. J. B., and Böhner, J.: Driving forces of mid-Holocene
 520 vegetation shifts on the upper Tibetan Plateau, with emphasis on changes in
 521 atmospheric CO₂ concentrations, <https://doi.org/10.1016/j.quascirev.2011.03.007>,
 522 *Quat. Sci. Rev.*, 30, 1907-1917, 2011.
- 523 Hou, A., Peng, S., and Zhou, G.: The study of the reactions of tree rings to the climate
 524 change and its applications, *Ecol. Sci.*, 18, 16-23,
 525 <https://doi.org/10.3969/j.issn.1008-8873.1999.03.004>, 1999 (in Chinese).
- 526 Hou, J., D'Andrea, W. J., Wang, M., He, Y., and Liang, J.: Influence of the Indian
 527 monsoon and the subtropical jet on climate change on the Tibetan Plateau since
 528 the late Pleistocene, *Quat. Sci. Rev.*, 163, 84-94,
 529 <https://doi.org/10.1016/j.quascirev.2017.03.013>, 2017a.
- 530 Hou, J., Huang, Y., Zhao, J., Liu, Z., Colman, S., and An, Z.: Large Holocene summer
 531 temperature oscillations and impact on the peopling of the northeastern Tibetan
 532 Plateau, *Geophys. Res. Lett.*, 43, 1323-1330,
 533 <https://doi.org/10.1002/2015GL067317>, 2016.
- 534 Hou, J., Tian, Q., Liang, J., Wang, M., and He, Y.: Climatic implications of hydrologic
 535 changes in two lake catchments on the central Tibetan Plateau since the last
 536 glacial, *J. Paleolimnol.*, 58, 257-273, <https://doi.org/10.1007/s10933-017-9976-9>,
 537 2017b.
- 538 Hu, G., Yu, L., Dong, Z., Jin, H., Luo, D., Wang, Y., and Lai, Z.: Holocene aeolian
 539 activity in the Headwater Region of the Yellow River, Northeast Tibet Plateau,
 540 China: A first approach by using OSL-dating, *Catena*, 149, 150-157,
 541 <https://doi.org/10.1016/j.catena.2016.09.014>, 2017.
- 542 Hu, G., Yu, L., Dong, Z., Lu, J., Li, J., Wang, Y., and Lai, Z.: Holocene aeolian
 543 activity in the Zoige Basin, northeastern Tibetan Plateau, China, *Catena*, 160,
 544 321-328, <https://doi.org/10.1016/j.catena.2017.10.005>, 2018.
- 545 Huang, E., Wang, P., Wang, Y., Yan, M., Tian, J., Li, S., and Ma, W.: Dole effect as a
 546 measurement of the low-latitude hydrological cycle over the past 800 ka, *Sci.*
 547 *Adv.*, 6, eaba4823, <https://doi.org/10.1126/sciadv.aba4823>, 2020.
- 548 Huo, B. Y., Guo, B. H., Liu, C. Y., Xu, H. M., and Jiang, Y. Q.: Significance of surface



- soil magnetic susceptibility in the Tengger Desert, *Arid Zone Res.*, 42, 613–621,
<https://doi.org/10.13866/j.azr.2025.04.04>, 2025 (in Chinese).
- Kasper, T., Haberzettl, T., Doberschütz, S., Daut, G., Wang, J., Zhu, L., Nowaczyk, N.
 R., and Mausbacher, R.: Indian ocean summer monsoon (IOSM)-dynamics
 within the past 4 ka recorded in the sediments of lake nam Co, central Tibetan
 plateau (China), *Quat. Sci. Rev.*, 39, 73–85,
<https://doi.org/10.1016/j.quascirev.2012.02.011>, 2012.
- Lehmkuhl, F., Schulte, P., Zhao, H., Hülle, D., Protze, J., and Stauch, G.: Timing and
 spatial distribution of loess and loess-like sediments in the mountain areas of the
 northeastern Tibetan Plateau, *Catena*, 117, 23–33,
<https://doi.org/10.1016/j.catena.2013.06.008>, 2014.
- Li, Y., Shang, H., Yan, S., Zhang, Z., Gao, M., Liu, S., Xi, T., Zhao, J., and Sun, W.:
 The Integrated Multi-proxy Paleoclimate Database for the Northeastern Tibetan
 Plateau, Mendeley Data, <https://doi.org/10.17632/cp7kftnb94.1>, 2026.
- Li, Y. and Morrill, C.: A holocene east asian winter monsoon record at the southern
 edge of the gobi desert and its comparison with a transient simulation, *Clim.
 Dynam.*, 45, 1219–1234, <https://doi.org/10.1007/s00382-014-2372-5>, 2015.
- Li, Y. and Peng, S.: Ecological and hydrologic evolution history in the sensitive zone
 of both east Asian summer monsoon and westerly since the last glacial maximum,
J. Mountain Sci., 20, 1266–1281, <https://doi.org/10.1007/s11629-022-7851-2>,
 2023.
- Li, Y., Peng, S., Hao, L., Zhou, X., and Li, H.: Interactions of the westerlies and Asian
 summer monsoon since the last deglaciation in the northeastern Qinghai-Tibet
 Plateau, *Paleoceanography and Paleoclimatology*, 37, e2022PA004548,
<https://doi.org/10.1029/2022PA004548>, 2022.
- Li, Y., Zhang, Z., Gao, M., Duan, J., Xue, Y., Shang, H., and Liu, S.: Monsoon
 marginal zone’s dry/wet status and their links with tropical ocean–atmosphere
 interactions during the holocene, *Sci. Bull.*,
<https://doi.org/10.1016/j.scib.2025.03.057>, 2025.
- Liang, E., Shao, X., Eckstein, D., and Liu, X.: Spatial variability of tree growth along
 a latitudinal transect in the Qilian Mountains, northeastern Tibetan Plateau, *Can.
 J. Forest Res.*, 40, 200–211, <https://doi.org/10.1139/X09-186>, 2010.
- Lister, G. S., Kelts, K., Zao, C. K., Yu, J.-Q., and Niessen, F.: Lake Qinghai, China:
 closed-basin lake levels and the oxygen isotope record for ostracoda since the
 latest Pleistocene, *Palaeogeogr. Palaeoclimatol. Palaeoecol.*, 84, 141–162,
[https://doi.org/10.1016/0031-0182\(91\)90041-O](https://doi.org/10.1016/0031-0182(91)90041-O), 1991.
- Liu, K.-B., Yao, Z., and Thompson, L. G.: A pollen record of Holocene climatic
 changes from the Dunde ice cap, Qinghai-Tibetan Plateau, *Geology*, 26, 135–138,
[https://doi.org/10.1130/0091-7613\(1998\)026<0135:APROHC>2.3.CO;2](https://doi.org/10.1130/0091-7613(1998)026<0135:APROHC>2.3.CO;2), 1998.
- Liu, W., Li, X., An, Z., Xu, L., and Zhang, Q.: Total organic carbon isotopes: a novel
 proxy of lake level from Lake Qinghai in the Qinghai-Tibet Plateau, China,
Chem. Geol., 347, 153–160, <https://doi.org/10.1016/j.chemgeo.2013.04.009>,
 2013.
- Liu, X., Dong, H., Yang, X., Herzsuh, U., Zhang, E., Stuut, J.-B. W., and Wang, Y.:



- 593 Late Holocene forcing of the Asian winter and summer monsoon as evidenced by
 594 proxy records from the northern Qinghai-Tibetan Plateau, *Earth Planet. Sci. Lett.*,
 595 280, 276-284, <https://doi.org/10.1016/j.epsl.2009.01.041>, 2009.
- 596 Liu, X., Herzschuh, U., Wang, Y., Kuhn, G., and Yu, Z.: Glacier fluctuations of
 597 Muztagh Ata and temperature changes during the late Holocene in westernmost
 598 Tibetan Plateau, based on glaciolacustrine sediment records, *Geophys. Res. Lett.*,
 599 41, 6265-6273, <https://doi.org/10.1002/2014GL060444>, 2014.
- 600 Liu, X., Shen, J., Wang, S., Wang, Y., and Liu, W.: Southwest monsoon changes
 601 indicated by oxygen isotope of ostracode shells from sediments in Qinghai Lake
 602 since the late Glacial, *Chin. Sci. Bull.*, 52, 539-544,
 603 <https://doi.org/10.1007/s11434-007-0086-3>, 2007.
- 604 Liu, X. D., Cheng, Z. G., Yan, L. B., et al.: Elevation dependency of recent and future
 605 minimum surface air temperature trends in the Tibetan Plateau and its
 606 surroundings, *Global Planet. Change*, 68, 164-174,
 607 <https://doi.org/10.1016/j.gloplacha.2009.03.017>, 2009a.
- 608 Liu, X. J., Lai, Z. P., Madsen, D. B., and Zeng, F. M.: Last deglacial and Holocene
 609 lake level variations of Qinghai Lake, north-eastern Qinghai-Tibetan plateau, *J.*
 610 *Quat. Sci.*, 30, 245-257, <https://doi.org/10.1002/jqs.2777>, 2015.
- 611 Liu, Y., Wang, Q., Ma, Z., Liu, X., and Yan, Y.: 540-year runoff reconstruction of
 612 upper Heihe River using tree-ring chronology, *J. Arid Land Resour. Environ.*, 23,
 613 93-97, <https://doi.org/10.13448/j.cnki.jalre.2009.03.011>, 2009b (in Chinese).
- 614 Lu, H., Wu, N., Liu, K.-B., Zhu, L., Yang, X., Yao, T., Wang, L., Li, Q., Liu, X., Shen,
 615 C., Li, X., Tong, G., and Jiang, H.: Modern pollen distributions in
 616 Qinghai-Tibetan Plateau and the development of transfer functions for
 617 reconstructing Holocene environmental changes, *Quat. Sci. Rev.*, 30, 947-966,
 618 <https://doi.org/10.1016/j.quascirev.2011.01.008>, 2011.
- 619 Long, H., Zhang, J., Zhang, S., Hou, Y., Wu, Y., Yang, N., Zhang, C., Cheng, L., Zhao,
 620 Z., Cheng, J., and Shen, J.: The westerlies-monsoon interaction shaped
 621 asymmetric lake expansions over the Tibetan Plateau in warming periods, *Sci.*
 622 *Bull.*, <https://doi.org/10.1016/j.scib.2025.07.023>, 2025.
- 623 Ma, Q., Zhu, L., Lü, X., Guo, Y., Ju, J., Wang, J., Wang, Y., and Tang, L.:
 624 Pollen-inferred Holocene vegetation and climate histories in Taro Co,
 625 southwestern Tibetan Plateau, *Chin. Sci. Bull.*, 59, 4101-4114,
 626 <https://doi.org/10.1007/s11434-014-0505-1>, 2014.
- 627 PAGES 2k Consortium: A global multiproxy database for temperature reconstructions
 628 of the Common Era, *Sci. Data*, 4, 170088, <https://doi.org/10.1038/sdata.2017.88>,
 629 2017.
- 630 Pan, B., Li, J., and Chen, F.: Qinghai-Tibetan Plateau: A driver and amplifier of global
 631 climate changes, *J. Lanzhou Univ. (Nat. Sci.)*, 31, 120-128,
 632 <https://doi.org/10.13885/j.issn.0455-2059.1995.03.024>, 1995 (in Chinese).
- 633 Qie, J.-Z., Zhang, Y., Trappmann, D., Zhao, Y.-H., Ballesteros-Canovas, J. A.,
 634 Favillier, A., and Stoffel, M.: Long-term reconstruction of flash floods in the
 635 Qilian Mountains, China, based on dendrogeomorphic methods, *J. Mt. Sci.*, 19,
 636 3163-3177, <https://doi.org/10.1007/s11629-022-7577-1>, 2022.



- 637 Qiu, J.: China: the third pole, *Nature*, 454, 393-396, <https://doi.org/10.1038/454393a>,
 638 2008.
- 639 Reimer, P. J., Austin, W. E. N., Bard, E., et al.: The IntCal20 Northern Hemisphere
 640 radiocarbon age calibration curve (0-55 cal kBP), *Radiocarbon*, 62, 725-757,
 641 <https://doi.org/10.1017/RDC.2020.41>, 2020.
- 642 Schiemann, R., Lüthi, D., and Schär, C.: Seasonality and interannual variability of the
 643 westerly jet in the Tibetan Plateau region, *J. Climate*, 22, 2940-2957,
 644 <https://doi.org/10.1175/2008JCLI2625.1>, 2009.
- 645 Shen, C., Liu, K. B., Morrill, C., Overpeck, J. T., Peng, J., and Tang, L.: Ecotone shift
 646 and major droughts during the mid-late Holocene in the central Tibetan Plateau,
 647 *Ecology*, 89, 1079-1088, <https://doi.org/10.1890/06-2016.1>, 2008.
- 648 Shi, F. F., Xiao, J. S., Li, L., Cao, X. Y., Li, X. D., Li, H. M.: A dataset of snow
 649 disaster grade monitoring based on meteorological elements in the pastoral area
 650 of the Tibetan Plateau, *China Scientific Data*, 7, 121-129,
 651 <https://doi.org/10.11922/11-6035.ncdc.2021.0029.zh>, 2022 (in Chinese).
- 652 Tang, L., Shen, C., Li, C., Peng, J., Liu, H., Liu, K.-B., Morrill, C., Overpeck, J. T.,
 653 Cole, J. E., and Yang, B.: Pollen-inferred vegetation and environmental changes
 654 in the central Tibetan Plateau since 8200 yr BP, *Sci. China Ser. D*, 52, 1104-1114,
 655 <https://doi.org/10.1007/s11430-009-0080-5>, 2009.
- 656 Tang, L., Shen, C., Liao, G., and Overpeck, J. T.: New high-resolution pollen records
 657 from two lakes in Xizang (Tibet), *Acta Bot. Sin.*, 41, 896-902, 1999.
- 658 Tang, L., Shen, C., Liu, K., and Overpeck, J. T.: Changes in South Asian monsoon:
 659 New high-resolution paleoclimatic records from Tibet, China, *Chin. Sci. Bull.*,
 660 45, 87-91, <https://doi.org/10.1007/BF02884911>, 2000.
- 661 Thompson, L. G., Yao, T., Davis, M. E., Henderson, K. A., Mosley-Thompson, E., Lin,
 662 P. N., Beer, J., Synal, H. A., Cole-Dai, J., and Bolzan, J. F.: Tropical climate
 663 instability: The last glacial cycle from a Qinghai-Tibetan ice core, *Science*, 276,
 664 1821-1825, <https://doi.org/10.1126/science.276.5320.1821>, 1997.
- 665 Tian, L., Masson-Delmotte, V., Stievenard, M., Yao, T., and Jouzel, J.: Tibetan Plateau
 666 summer monsoon northward extent revealed by measurements of water stable
 667 isotopes, *J. Geophys. Res.-Atmos.*, 106, 28081-28088,
 668 <https://doi.org/10.1029/2001JD900186>, 2001.
- 669 Wang, F. F., Hou, G. L., Hou, X. Q., and Wang, Q. B.: Researches on Climate
 670 Parameters for Quantitative Reconstruction Based on Surface Pollen in the
 671 Tibetan Plateau, *Acta Geol. Sin.*, 38, 529-536,
 672 <https://doi.org/10.3975/cagsb.2017.04.10>, 2017 (in Chinese).
- 673 Wang, H., Dong, H., Zhang, C. L., Jiang, H., Zhao, M., Liu, Z., Lai, Z., and Liu, W.:
 674 Water depth affecting thaumarchaeol production in Lake Qinghai, northeastern
 675 Qinghai-Tibetan plateau: Implications for paleo lake levels and paleoclimate,
 676 *Chem. Geol.*, 368, 76-84, <https://doi.org/10.1016/j.chemgeo.2014.01.009>, 2014.
- 677 Wang, N., Yao, T., Pu, J., Zhang, Y., and Sun, W.: Climatic and environmental
 678 changes during the past millennium recorded by the Malan ice core in the
 679 northern Tibetan Plateau, *Sci. China Ser. D*, 36, 723-732,
 680 <https://doi.org/10.3321/j.issn:1006-9267.2006.08.005>, 2006 (in Chinese).



- 681 Wang, S. Y., Wang, Q. L., Kang, S. C., Ding, Y. J., Xu, Y. A., and Huang, X. B.:
 682 Characteristics and controlling factors of stable isotopes in precipitation based on
 683 9-year monitoring data in the monsoon transition zone of the Tibetan Plateau,
 684 Agric. For. Meteorol., <https://doi.org/10.1016/j.agrformet.2025.110889>, 2026.
- 685 Wu, D., Zhou, A., Zhang, J., Chen, J., Li, G., Wang, Q., Chen, L., Madsen, D., Abbott,
 686 M., Cheng, B., and Chen, F.: Temperature-induced dry climate in basins in the
 687 northeastern Tibetan Plateau during the early to middle Holocene, Quat. Sci.
 688 Rev., 237, 106311, <https://doi.org/10.1016/j.quascirev.2020.106311>, 2020.
- 689 Wu, G. X., Liu, Y. M., Liu, X., et al.: How the heating over the Tibetan Plateau affects
 690 the Asian climate in summer, Chin. J. Atmos. Sci., 29, 47-56, 2005 (in Chinese).
- 691 Wünnemann, B., Demske, D., Tarasov, P. E., Kotlia, B. S., Reinhardt, C., Bloemendal,
 692 J., Diekmann, B., Hartmann, K., Krois, J., and Riedel, F.: Hydrological evolution
 693 during the last 15 kyr in the Tso Kar lake basin (Ladakh, India), derived from
 694 geomorphological, sedimentological and palynological records, Quat. Sci. Rev.,
 695 29, 1138-1155, <https://doi.org/10.1016/j.quascirev.2010.02.017>, 2010.
- 696 Yang, B., He, M., Melvin, T. M., Osborn, T. J., and Briffa, K. R.: Climate control on
 697 tree growth at the upper and lower treelines: a case study in the Qilian Mountains,
 698 Tibetan Plateau, PLoS One, 8, e69065,
 699 <https://doi.org/10.1371/journal.pone.0069065>, 2013.
- 700 Yang, B., Qin, C., Shi, F., and Sonechkin, D. M.: Tree ring-based annual streamflow
 701 reconstruction for the Heihe River in arid northwestern China from AD 575 and
 702 its implications for water resource management, Holocene, 22, 773-784,
 703 <https://doi.org/10.1177/0959683611430411>, 2012.
- 704 Yang, B., Qin, C., Wang, J., He, M., Melvin, T. M., Osborn, T. J., and Briffa, K. R.: A
 705 3,500-year tree-ring record of annual precipitation on the northeastern Tibetan
 706 Plateau, Proc. Natl. Acad. Sci., 111, 2903-2908,
 707 <https://doi.org/10.1073/pnas.1319238111>, 2014.
- 708 Yao, T., Thompson, L. G., Shi, Y., Qin, D., Jiao, K., Yang, Z., Tian, L., and Thompson,
 709 E. M.: Climate variation since the Last Interglaciation recorded in the Guliya ice
 710 core, Sci. China Ser. D, 40, 662-668, <https://doi.org/10.1007/BF02877697>, 1997.
- 711 Yin, Z.-Y., Zhu, H., Huang, L., and Shao, X.: Reconstruction of biological drought
 712 conditions during the past 2847 years in an alpine environment of the
 713 northeastern Tibetan Plateau, China, and possible linkages to solar forcing,
 714 Global Planet. Change, 143, 214-227,
 715 <https://doi.org/10.1016/j.gloplacha.2016.04.010>, 2016.
- 716 Zan, J. B., Ning, W. X., Yang, S. L., Fang, X. M., Kang, J., and Luo, Y. L.: Magnetic
 717 Variations in Surface Soils from the Tibetan Plateau and Its Adjacent Regions:
 718 Implications for Delineating the Climatic Boundary, Adv. Earth Sci., 37, 14-25,
 719 <https://doi.org/10.11867/j.issn.1001-8166.2021.125>, 2022 (in Chinese).
- 720 Zhan, S., Wang, K., Zhang, L., et al.: Species-specific growth responses to climatic
 721 factors in the eastern Qilian Mountains, Chin. J. Ecol., 38, 2007-2014,
 722 <https://doi.org/10.13292/j.1000-4890.201907.021>, 2019 (in Chinese).
- 723 Zhang, P. X., Zhang, B. Z., Qian, G. M., Li, H. J., and Xu, L. M.: The study of
 724 paleoclimatic parameter of Qinghai Lake since Holocene, Quat. Sci., 3, 225-236,



725 1994 (in Chinese).
726 Zhang, Y., Shao, X. M., Yin, Z.-Y., and Wang, Y.: Millennial minimum temperature
727 variations in the Qilian Mountains, China: evidence from tree rings, *Clim. Past*,
728 10, 1763-1778, <https://doi.org/10.5194/cp-10-1763-2014>, 2014.