



# A pan-tropical 5-km monthly L-band vegetation optical depth dataset from pan-sharpening-based downscaling

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**Abstract.** L-band Vegetation optical depth (L-VOD), as a microwave-derived vegetation indicator, has been widely applied in the monitoring of vegetation dynamics. However, the spatial resolution of 25-km or coarser in existing L-band VOD products limits their applications in ecological monitoring requiring a higher level of spatial details. To mitigate this limitation, we introduce a pan-sharpening-based downscaling method to improve the spatial resolution of L-VOD. By fusing the spatial structural features of the aggregated 5-km resolution European Space Agency Climate Change Initiative (ESA CCI) aboveground biomass (AGB) product, the SMOS L-VOD product over tropical regions was downscaled to generate a monthly 5-km resolution L-VOD dataset spanning 2015 to 2021. The downscaling model demonstrated high accuracy, with a correlation coefficient ( $R$ ) of 0.95 and a root mean square error ( $RMSE$ ) of 0.11 when comparing the simulated 25-km L-VOD ( $L - VOD_{25km}^{sim}$ ) with the original L-VOD ( $L - VOD_{25km}$ ) product. Spatially, the 5-km resolution L-VOD ( $L - VOD_{5km}$ ) yielded a strong correlation with above-ground biomass ( $R = 0.91$ ,  $R^2 = 0.86$ ), and temporally dynamics, it accurately characterized the LAI variations of short vegetation and forest area loss at the pixel level over the study period. The results demonstrate that our downscaling method can effectively enhance the spatial resolution of L-VOD while preserving its original spatiotemporal dynamics, and is capable of capturing forest disturbance. This dataset can be downloaded at <https://doi.org/10.11888/RemoteSen.tpd.303391> (Shi and Fan, 2026).

35 **Keywords:** L-band Vegetation Optical Depth (L-VOD), Downscaling, Pan-sharpening, Pan-tropical, 5-km resolution



## 1 Introduction

Vegetation optical depth (VOD) provides critical insights into the dynamics of the global carbon in the context of changing environmental and climatic conditions (Skulovich et al., 2024). It provides reliable estimates of vegetation biomass (Yang et al., 2023; Wigneron et al., 2024; Zhang et al., 2025; Tong et al., 2020) and vegetation water content (VWC) (Felton et al., 2025; Jackson and Schmugge, 1991; Piles et al., 2017; Tong et al., 2018), based on the parameterization of vegetation attenuation and emission (Wigneron et al., 2017). Global long-term time-series VOD products can be obtained from various active and passive observations at different microwave frequency bands, such as C-band (Liu et al., 2021; Liu et al., 2023; Owe et al., 2008), X-band (Karthikeyan et al., 2019; Wang et al., 2021), and L-band (Chaubell et al., 2021; Konings et al., 2016; Wigneron et al., 2021). L-band exhibits greater penetration capability through the vegetation canopy than C-band and X-band, and thus shows a considerably lower saturation effect compared with other frequencies (Xing et al., 2024), particularly where vegetation canopy is dense and biomass density is high (Brandt et al., 2018). Therefore, L-VOD provides a more reliable means of monitoring forest structural properties and their temporal dynamics in high-biomass regions. Two currently operational L-band missions – Soil Moisture and Ocean Salinity (SMOS) (Kerr et al., 2010) and Soil Moisture Active Passive (SMAP) (Entekhabi et al., 2008; Xing et al., 2024) – have shown a strong capability for monitoring both the vegetation and soil dynamics globally (Wigneron et al., 2017). Many L-VOD datasets have been generated in recent years, offering reliable and dependable long-term L-band VOD time series for use in ecological and climate-related analyses. A range of distinct datasets are known as SMOS-IC (Fernandez-Moran et al., 2017), SMAP-IB (Li et al., 2022a; Xiao et al., 2024), SMOSMAP-IB (Li et al., 2022b), SMAPMT-DCA (Konings et al., 2017) and VODCA v2 (Zotta et al., 2024).

Although satellite-based VOD retrievals provide essential regional to continental-scale forest biomass information at daily temporal resolution, their coarse spatial resolution ( $\geq 25$ -km) limits their application to accurately monitor local, fine-scale, or patchy forest biomass dynamics. At such coarse resolution, each grid cell typically encompasses multiple land cover types, which prevents investigations into the temporal dynamics of diverse vegetation categories (Vreugdenhil et al., 2020). Moreover, for each VOD pixel, the properties of the microwave signal are significantly influenced the following factors: the density of the vegetation, the roughness features of the soil surface, together with the VWC (Piles et al., 2009; Zhou et al., 2022), which leads to exceptionally strong spatial variability of VOD within individual pixels (Owe et al., 2002). Additionally, higher-resolution VOD serves as a critical parameter in microwave radiative transfer models for soil moisture (SM) retrievals, while further contributing prominently to higher-resolution SM inversion (Vreugdenhil et al., 2020). Thus, higher-resolution VOD is becoming increasingly crucial for vegetation dynamics monitoring and hydrological studies at regional and local level. Additionally, it offers the possibility of filling gaps in optical vegetation indices (VIs) that are unavailable in the presence of cloud cover.

Researchers have employed both passive and active microwave radiation to retrieve higher-resolution VOD. For passive retrieval, the Land Parameter Retrieval Model (LPRM) has been applied to sharpened AMSR-E brightness temperatures, yielding soil moisture and VOD at an enhanced resolution of approximately 10-km across Australia (Gevaert et al., 2016).



Active microwave sensors and especially SARs are capable of offering far higher spatial resolutions than passive radiometers  
70 (Peng et al., 2017; Liu et al., 2023). C-band backscatter data acquired by Sentinel-1 have been utilized to retrieve VOD products  
with a spatial resolution of 1-km (El Hajj et al., 2019; Zhou et al., 2022). However, retrieving VOD from SAR is often difficult  
as a result of the integrated influences of canopy architecture, surface roughness, and moisture content on SAR backscatter  
(Wagner et al., 2007; Frappart et al., 2020).

Statistical downscaling, which is also referred to as multi-parameter fusion-based downscaling (Jing et al., 2022), introduces  
75 higher resolution auxiliary parameters and fuses them with the coarse-resolution VOD data to obtain the finer resolution VOD  
data. The core of these methods lies in developing a statistical model of the VOD-auxiliary variable relationship at a low  
resolution. This model is subsequently applied to generate high-resolution estimates, based on the key assumption that the  
relationship does not vary with spatial scale (Jing et al., 2022). In terms of auxiliary data sources, in one respect, scholars have  
utilized active microwave auxiliary data (Jiang et al., 2024) and proposed VOD downscaling methods using radar backscatter  
80 coefficients (Bousquet et al., 2019). In another respect, investigators have applied optical or thermal infrared auxiliary data  
(Jiang et al., 2024) and developed downscaling approaches to estimate VOD, such as based on vegetation indices NDVI (Xinyu  
et al., 2025). The conventional statistical approaches discussed above are subject to the limitation of the assumption of linearity  
between VOD and auxiliary variables, consequently failing to capture their complex nonlinear interactions.

Previous studies have utilized random forest (RF) models to downscale VOD with multi-source auxiliary information. The  
85 downscaling of VOD using RF firstly builds up a relationship between low-resolution VOD and higher-resolution vegetation  
indices (e.g., NDVI, LAI), and subsequently the relationship is used to convert higher-resolution vegetation indices into higher-  
resolution VOD. This means that the estimated higher-resolution VOD can be considered to be a conversion of vegetation  
indices values. For instance, previous studies (Mohite et al., 2022; Yan et al., 2025) used RF to downscale VOD, by building  
up the relationship between VOD and LST, NDVI, DEM.

90 In the aforementioned VOD downscaling methods, the auxiliary higher-resolution variables include optical vegetation indices  
like NDVI (Mohite et al., 2022; Xinyu et al., 2025; Yan et al., 2025) and radar data (e.g., L-band backscattering coefficients  
from the ALOS-1 mission) (Bousquet et al., 2019). NDVI exhibits significant saturation phenomena in regions with high  
vegetation coverage (Xinyu et al., 2025), while L-band radar backscatter coefficients (HV polarization) typically demonstrate  
signal saturation in evergreen broadleaf forests when exceeding 0.07 (Bousquet et al., 2019). Although L-VOD and AGB  
95 exhibit an obvious non-linear saturating relationship, the degree of saturation is relatively weak (Li et al., 2021). Furthermore,  
the Climate Change Initiative (CCI) AGB product provides 100 m resolution data (derived from multi-source fusion and spatial  
resampling (Boitard et al., 2024), making CCI-AGB a suitable high-resolution auxiliary variable for image fusion-based  
downscaling methods to enhance spatial detail representation.

An alternative downscaling method based on the morphological filtering-based multiresolution analysis (MF-MRA) technique  
100 (Restaino et al., 2016) uses 5-km resolution AGB data to sharpen the low-resolution VOD data. This method utilizes the VOD  
as a baseline, employing morphological operators to extract spatial structural features—such as textural and edge



information—from the AGB data, and injects them into the VOD. Consequently, the method could improve spatial resolution through structural feature injection, without relying on the absolute values of auxiliary data.

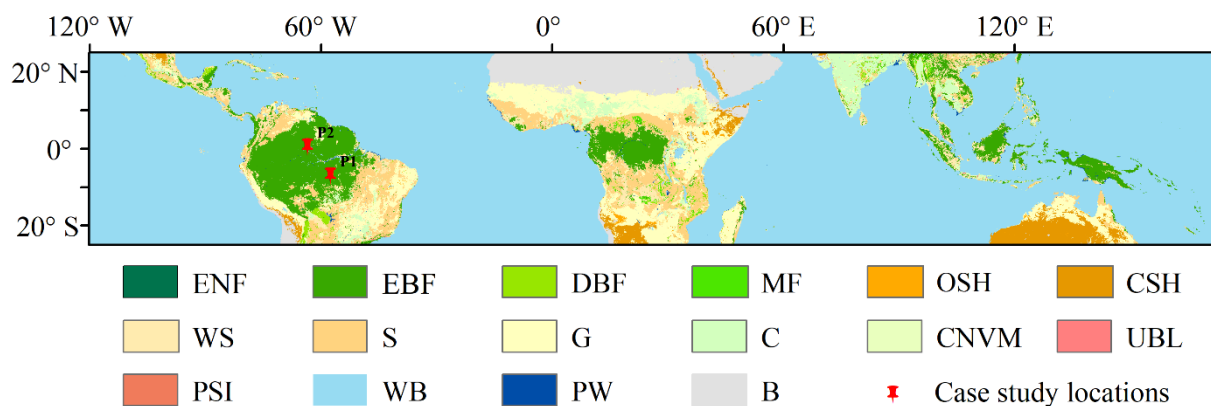
105 Within the present work, we adopt the MF-MRA fusion approach commonly used in panchromatic sharpening to enhance the spatial resolution of L-VOD. By fusing the monthly 25-km resolution L-VOD data with the concurrent monthly 5-km resolution aggregated ESA CCI AGB (original 100-m resolution) from 2015 to 2021, a monthly 5-km resolution L-VOD dataset for the period 2015–2021 is generated. This strategy not only improves the spatial resolution of the L-VOD product but also effectively preserves the potential nonlinear structural information between L-VOD and AGB in the fusion process. The primary aims of this work are as follows: (1) to introduce a novel L-VOD downscaling framework, which is based on  
110 MF-MRA to enhance the spatial resolution of SMOS-IC L-VOD from 25-km to 5-km, (2) to demonstrate the applicability and robustness of the proposed downscaling approach, and (3) to validate the spatiotemporal consistency of the downscaled L-VOD in regards to temporal variation and spatial distribution. This was done through comparisons with AGB, LAI, and forest loss datasets.

## 2 Datasets

### 115 2.1 Study area and land cover

Our study focused on the tropical region spanning 25° N to 25° S (Fig. 1), covering tropical America, Africa, and Asia. This region contains the world’s major rainforests, which are characterized by greater rainfall and denser vegetation than other regions (Ma et al., 2023). These forests are dynamic ecosystems that undergo ecological succession, natural perturbations, deforestation, and land degradation (Abbas et al., 2020), and provide more than half the uptake of terrestrial atmospheric  
120 carbon dioxide ( $CO_2$ ) (Stephens et al., 2007; Lewis et al., 2009). The spatial heterogeneity in both climate and vegetation types makes this region ecologically representative and suitable for a large-scale study on downscaling L-VOD.

Land cover data for this study were derived from the MODIS MCD12C1 dataset (Friedl et al., 2010), which offers 5-km resolution land cover classifications for 2019 following the International Geosphere-Biosphere Programme (IGBP) scheme. We performed an aggregation at a 25-km grid resolution using the predominant category. “Predominant” denotes the class  
125 with the highest frequency of 5-km pixels within each 25-km grid.



130 **Figure 1.** The land cover categories for 2019 over the tropical study domain are obtained from the MODIS dataset at 5-km spatial resolution. Deciduous needleleaf forest (DNF), Evergreen broadleaf forest (EBF), Evergreen needleleaf forest (ENF), Mixed forests (MF), Closed Shrublands (CSH), Open Shrublands (OSH), Woody Savannas (WS), Savannas (S), Grasslands (G), Permanent Wetlands (PW), Croplands (C), Cropland/Natural Vegetation Mosaics (CNVM), Barren (B). The two detailed study regions (P1 and P2), each comprising a 2×2 group of Soil Moisture and Ocean Salinity (SMOS) grid cells (P1: central latitude, 7.0° S; longitude, 57.5° W; P2: central latitude, 0.53° N; longitude, 63.43° W) are shown in Fig. 6 and Fig. 10.

## 2.2 L-VOD product

The L-VOD was retrieved from SMOS satellite descending and ascending orbit observations via the SMOS-IC algorithm, covering the period from 2010–2021 (Li et al., 2020). Radio frequency interference (RFI) can distort the signals received by satellite sensors, potentially introducing uncertainty into the VOD retrieval process. L-VOD data affected by RFI were excluded utilizing the methodology described by Wigneron et al. (2021), which applies a threshold of 8 K to the root mean square error (*RMSE*) between measured and simulated brightness temperature. Furthermore, bilinear interpolation, commonly used in VOD processing (Li et al., 2021; Liu et al., 2018), was applied to reproject and resample the L-VOD dataset to a 25-km grid. For this study, L-VOD data for the period 2015–2021 were utilized.

## 2.3 AGB

AGB estimates in the ESA CCI Biomass maps (Santoro and Cartus, 2024) are based on radar backscatter intensity measurements. According to Boitard et al. (2024), these measurements were obtained from the PALSAR2 sensor onboard ALOS2 and the Sentinel-1 satellites, which were developed under the ESA CCI Biomass project. LiDAR metrics and surface reflectance were also used in the estimation procedure. These AGB maps are updated on an annual basis and have a spatial resolution of 100 x 100m. This work uses version 5.01 of the maps for the period 2015–2021. The original 100-m resolution AGB maps were aggregated to a 5-km resolution using mean values to serve as input for the downscaling model.

## 2.4 MODIS Vegetation Indices

NDVI and LAI, which are frequently adopted as proxies for green vegetation coverage (Grant et al., 2016; Lawrence et al., 2014; Liu et al., 2011; Liu et al., 2021) were used to assess the performance of downscaled L-VOD. The MOD13A1 Version



6.0 dataset provides 16-day composite NDVI records at a spatial resolution of 500-m, derived by selecting and averaging the highest-quality observations during each 16-day compositing period (Didan, 2015). The MODIS MOD15A2H Version 6 dataset provides 8-day averaged LAI values on a 500-m spatial grid (Myneni et al., 2015). Over the study period, high-quality NDVI and LAI pixels flagged by the quality control layers were chosen and resampled to a 5-km spatial grid using an averaging method.

## 2.5 Forest Cover Loss

Hansen et al. (2013) developed a forest cover loss dataset (version 1.11) that maps annual forest loss at a 30-m spatial grid for the 2000–2023 period. In this product, forest cover loss is defined as temporary or permanent transition from forest to nonforest. This dataset includes a “lossyear” layer that denotes the year of detected forest cover loss, covering all forest loss events observed from 2000 to 2023 (Fan et al., 2022). Here, annual forest percentage loss rates over the period 2015–2021 were computed based on the resolution of downscaled L-VOD grid. Specifically, we quantified the proportion of the total forest loss areas as identified in the “lossyear” within each downscaled L-VOD grid cell (~5-km). Of note, this forest loss product represents temporary or permanent tree cover loss resulting from disturbances such as wildfires or clearcutting.

## 2.6 Canopy Height

We used a global forest canopy height (CH) product at 30-m resolution for 2019, produced by the Global Land Analysis and Discovery team at the University of Maryland (UMD GLAD) (Potapov et al., 2021), which provides seamless coverage from 52°N to 52°S latitudes. To align with the spatial scale employed in this research, the original dataset was aggregated to 5-km by averaging.

## 3 Methods

Fig. 2 presents the flow chart of our sequential downscaling process, consisting of three steps: (1) data preparation, (2) downscaling, and (3) evaluation.

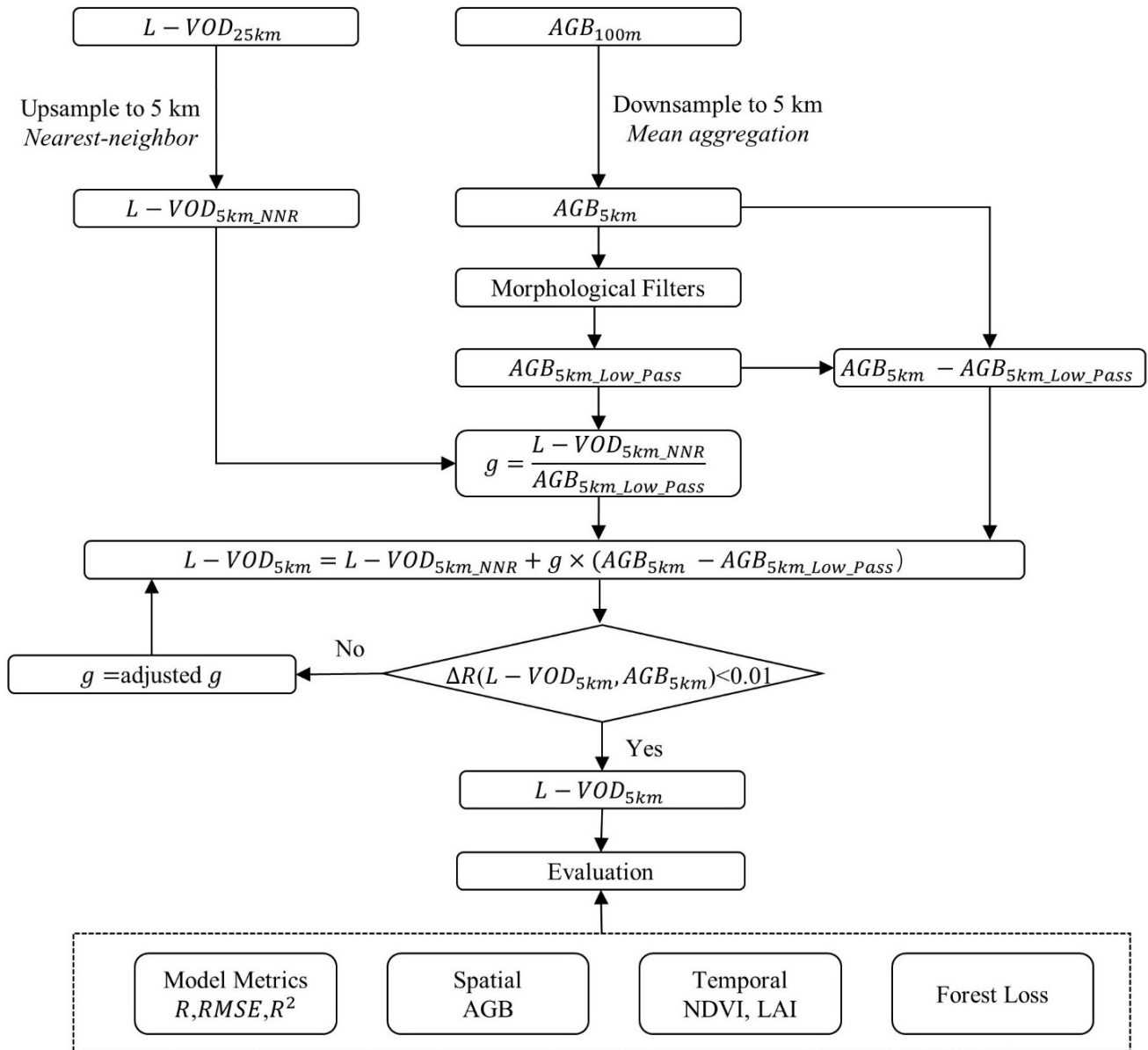


Figure 2. Flowchart illustrating the workflow for generating and validating downscaled L-VOD products.

### 3.1 Data preparation

175 We adopted the 25-km resolution L-VOD and the 100-m resolution AGB as inputs, namely  $L - VOD_{25km}$  and  $AGB_{100m}$  in Fig. 2, respectively. The AGB data provided higher-resolution spatial details to guide the enhancement of L-VOD resolution. The monthly  $L - VOD_{25km}$  data from 2015 to 2021 were first combined into 12-band images for each year. The L-VOD were then resampled to 5-km resolution ( $L - VOD_{5km\_NNR}$ ) through the nearest-neighbor resampling (NNR) approach, which avoids



interpolation-based smoothing and preserves the fidelity of the original data (Parker et al., 2007; Porwal and Katiyar,  
180 2014).  $AGB_{100m}$  data were aggregated to 5-km resolution ( $AGB_{5km}$ ) through arithmetic averaging.

### 3.2 The L-VOD downscaling method

The 5-km resolution L-VOD is estimated from the  $L - VOD_{5km\_NNR}$  by injecting spatial details from the  $AGB_{5km}$  using the MF-MRA method by the following equation:

$$L - VOD_{5km} = L - VOD_{5km\_NNR} + g \times (AGB_{5km} - AGB_{5km\_Low\_Pass}) \quad (1)$$

where  $L - VOD_{5km\_NNR}$  stands for the monthly L-VOD data at 5-km resolution using NNR method. To extract the low-  
185 frequency component of the AGB data ( $AGB_{5km\_Low\_Pass}$ ), morphological filtering was applied to the 5-km resolution AGB  
( $AGB_{5km}$ ). In this study, we assume that both  $AGB_{5km}$  and ( $AGB_{5km\_Low\_Pass}$ ) remain constant within each year. Thus, for a  
given year, the same annual AGB value is applied to all twelve months in the Eq. (1) to match the monthly L-VOD data.  
Spatial details in AGB ( $AGB_{5km} - AGB_{5km\_Low\_Pass}$ ) were extracted and injected by  $g$  (see Eq. (2)), and subsequently added  
to the monthly  $L - VOD_{5km\_NNR}$  to enhance its spatial resolution. As a result, the downscaled  $L - VOD_{5km}$  is a monthly  
190 dataset.

The injection gains  $g$  in Eq. (1) is a pixel-wise injection gain, representing the contribution of annual AGB to each monthly  
L-VOD. A total of 84 gain maps (7 years  $\times$  12months) were obtained to downscale the 5-km resolution monthly L-VOD series  
over the 7-year period (2015–2021), with each monthly gain map computed based on the corresponding annual AGB, as  
described in Eq. (2).

$$g = \frac{L - VOD_{5km\_NNR}}{AGB_{5km\_Low\_Pass}} \quad (2)$$

195 Note that injection gains  $g$  contain uncertainties, because the  $L - VOD_{5km\_NNR}$  is not the true 5-km L-VOD although  
 $L - VOD_{5km\_NNR}$  could represent the spatial-temporal information of  $L - VOD_{25km}$ . To reduce the uncertainties of  $g$  and  
enhance the accuracy and robustness of  $L - VOD_{5km}$ , the injection gains  $g$  were iteratively optimized by maximizing the  
Pearson correlation coefficient ( $R$ ) between the  $L - VOD_{5km}$  and the  $AGB_{5km}$ , formally:

$$g = \text{arg}R(L - VOD_{5km}, AGB_{5km}) \quad (3)$$

in each iteration, the updated  $L - VOD_{5km}$  was recalculated using the new  $g$ , and the correlation with  $AGB_{5km}$  was re-  
200 evaluated. In the initial iteration, the maximum value of  $g$  across all pixels and all layers was extracted and defined as the  
initial  $g_{limit}$ . In each subsequent iteration, if a pixel's  $g$  exceeded the current  $g_{limit}$ , it was set equal to  $g_{limit}$ , i.e.,  $g =$   
( $g, g_{limit}$ ). After each iteration,  $g_{limit}$  was reduced by a fixed step of 0.01, i.e.,  $g_{limit} = g_{limit} - 0.01$ . Iterations were  
continued until  $\Delta R < 0.01$  ( $\Delta R$  is defined as the difference in  $R$  between iterations). Formally, the stopping criterion is:

$$\Delta R = |R^k - R^{k-1}| < 0.01 \quad (4)$$

where  $R^k$  and  $R^{k-1}$  are the correlation coefficients at the  $k$ -th and  $(k - 1)$ -th iterations, respectively, and  $R$  is correlation  
205 coefficient between the  $L - VOD_{5km}$  and the  $AGB_{5km}$ .



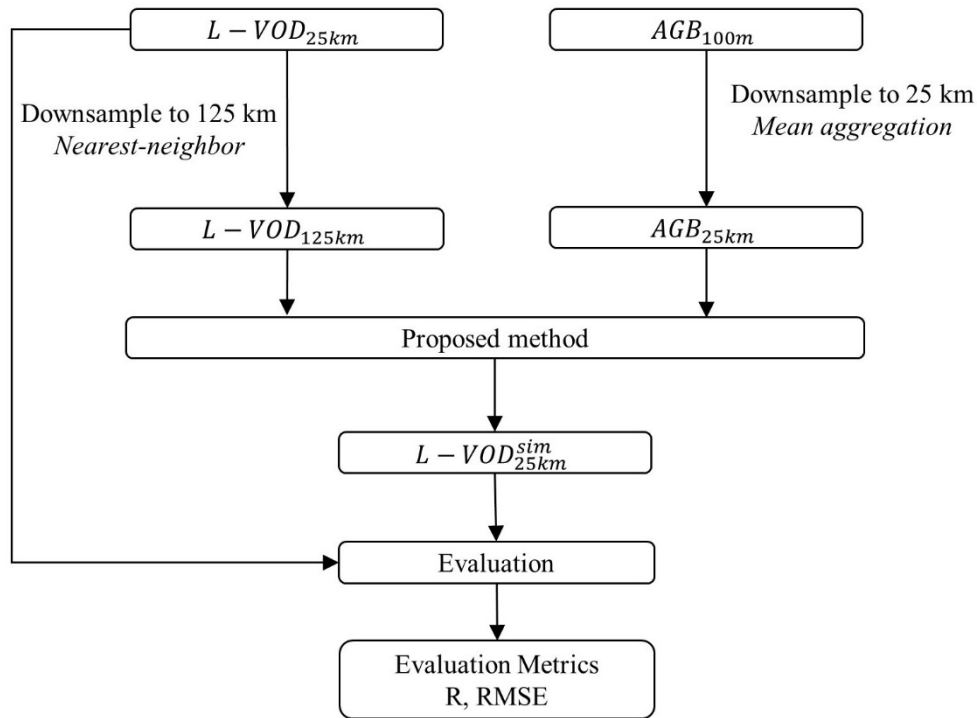
To illustrate the proposed method, we take the year 2017 as an example. The 2017 AGB data at 5-km resolution and its low-frequency component obtained by morphological filtering are assumed constant for all twelve months of 2017. Thus, for each month  $m = 1, 2, 3, \dots, 12$ , the downsampled L-VOD at 5-km resolution is calculated as:

$$L - VOD_{5km}^{2017,m} = L - VOD_{5km\_NNR}^{2017,m} + g^{2017,m} \times (AGB_{5km}^{2017,m} - AGB_{5km\_Low\_Pass}^{2017,m}) \quad (5)$$

where  $L - VOD_{5km}^{2017,m}$  and  $L - VOD_{5km\_NNR}^{2017,m}$  represent the downsampled and nearest neighbor resampled monthly L-VOD for month  $m$ , respectively.  $g^{2017,m}$  is the injection gain map corresponding to month  $m$ .  $AGB_{5km}^{2017,m}$  is the annual AGB at 5-km resolution; and  $AGB_{5km\_Low\_Pass}^{2017,m}$  is its low-frequency component derived by morphological filtering.

### 3.3 Accuracy assessment

Our analysis involves two main aspects, namely the methodological accuracy of the downscaling and the evaluation of the 5-km product.



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**Figure 3. Schematic diagram of the validation approach. Methodological evaluation at 25-km resolution.**

(i) Evaluation of the downscaling method. The approach for verifying the downscaling method is illustrated in Fig. 3. First, the original 25-km L-VOD (denoted as  $L - VOD_{25km}$ ) is aggregated to 125-km spatial resolution via the nearest-neighbor algorithm (denoted as  $L - VOD_{125km}$ ). In parallel,  $AGB_{25km}$  data is generated from the original  $AGB_{1km}$  by mean aggregation.



220 Subsequently, both  $L - VOD_{125km}$  and  $AGB_{25km}$  are utilized as inputs for the proposed method to generate a downscaled 25-  
km L-VOD, denoted as  $L - VOD_{25km}^{sim}$ . Thus, the original  $L - VOD_{25km}$  acts as the reference benchmark, while  $L - VOD_{25km}^{sim}$   
represents the output of the downscaling method. Method performance was evaluated by computing the R and RMSE obtained  
from the comparison between the  $L - VOD_{25km}^{sim}$  and  $L - VOD_{25km}$ . This validation framework has been adopted in previous  
studies and proven to be a reliable method for assessing core model performance, as documented by Merlin et al. (2008). The  
225 equations corresponding to these metrics are listed as follows:

$$R = \frac{Cov(L - VOD_{25km}^{sim}, L - VOD_{25km})}{\sigma(L - VOD_{25km}^{sim})\sigma(L - VOD_{25km})} \quad (6)$$

$$RMSE = \sqrt{\frac{\sum(L - VOD_{25km}^{sim} - L - VOD_{25km})^2}{N}} \quad (7)$$

where  $Cov$  is the covariance and  $\sigma$  is the standard deviation.  $N$  is the sample size. Statistical significance was set at  $p < 0.05$ .

(ii) Evaluation of  $L - VOD_{5km}$  product. Regarding L-VOD datasets, no universal agreement has been reached on suitable  
reference VOD estimates from field observations or simulation models for validation purposes (Li et al., 2020). Hence, no  
single evaluation procedure is universally applicable, and instead diverse assessment strategies can be implemented under  
230 specific conditions. These conditions may include different vegetation regimes, such as grasslands and savannas (low-to-  
moderate cover) or tropical rainforests and boreal forests (dense cover), as well as specific sites affected by disturbances like  
deforestation and fire events (Grant et al., 2016; Tian et al., 2016; Xing et al., 2024).

The spatial correlation between  $L - VOD_{5km}$  and AGB (Li et al., 2021) was used to assess the robustness of  $L - VOD_{5km}$   
datasets. It should be noted that the AGB dataset adopted here is the same as that used in the downscaling process, which  
235 provides spatial structural information for the downscaling of L-VOD. Therefore, this evaluation reflects spatial consistency  
between the  $L - VOD_{5km}$  and  $AGB_{5km}$  rather than an independent accuracy validation. As a parameter characterizing  
microwave signal attenuation by vegetation, VOD displays spatial patterns across frequencies that closely resemble those of  
the IGBP global land cover classification (Konings et al., 2017). A set of VIs, including LAI and NDVI, were selected to be  
compared with  $L - VOD_{5km}$ , since they are commonly utilized as proxies for green vegetation coverage (Weber et al., 2020).  
240 Previous studies have extensively validated the representativeness of VOD by cross-comparing it with optical vegetation  
indices. (Grant et al., 2016; Lawrence et al., 2014; Li et al., 2020). Temporal correlations at the pixel level were calculated  
between  $L - VOD_{5km}$  and these VIs.  $L - VOD_{5km}$  spatiotemporal consistency with forest loss was examined, since Qin et al.  
(2021) observed a strong correlation between L-VOD dynamics and variations in forest loss within the Amazon Basin.

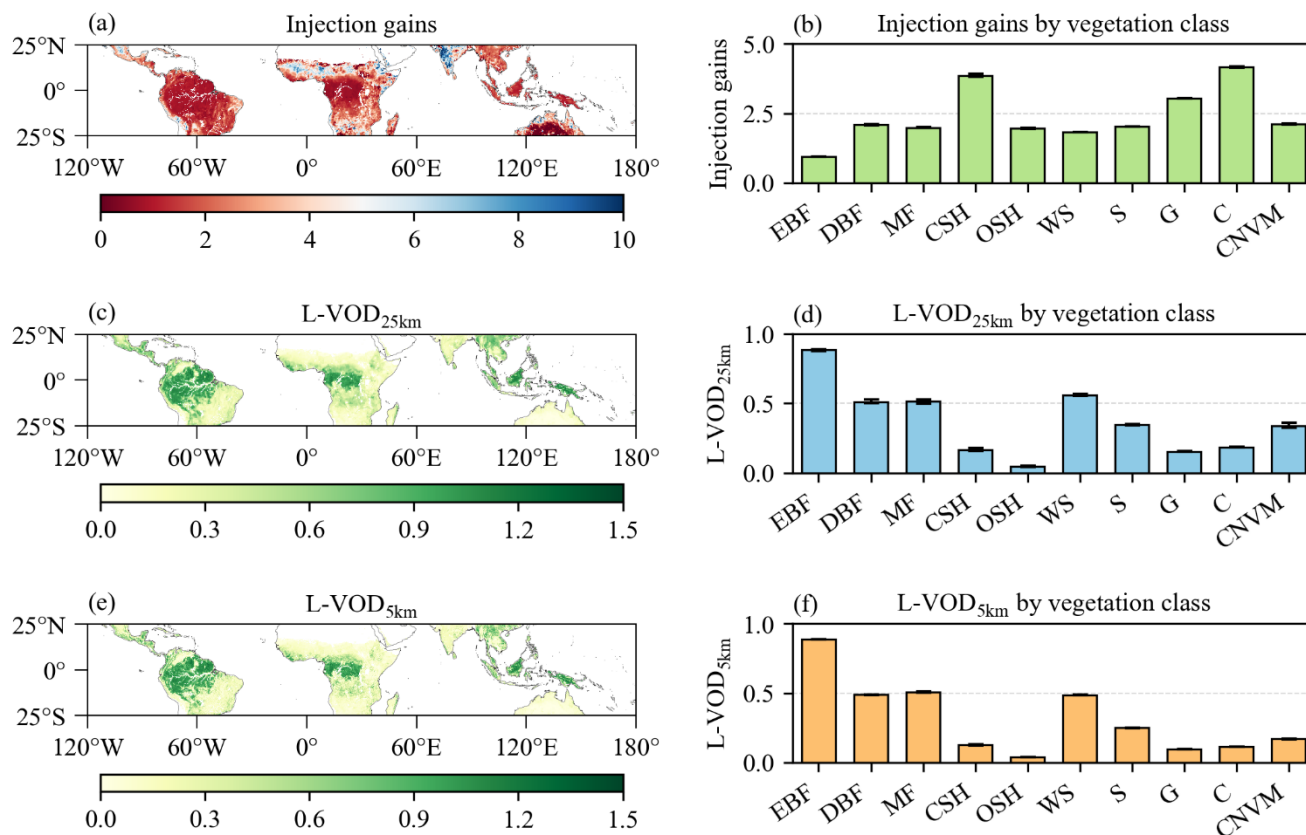


## 4 Results

### 245 4.1 Spatial pattern of the downscaled L-VOD data

In the process of downscaling, the spatial details being injected (i.e., the injection gains  $g$  defined in Eq. (1)) exhibit distinct patterns across various vegetation classes (Fig. 4a). In regions of tall, homogeneous vegetation, such as the Amazon,  $g$  is low. Our observation is in line with the saturation behavior of L-band signals, which occurs in dense vegetation canopies and at elevated values of above-ground biomass. In contrast, high  $g$  values appeared in zones dominated by short or heterogeneous  
250 vegetation, such as the Sahel. This pattern is consistent with the fact that the L-band signal has not yet reached saturation and is more sensitive to variations in AGB (Li et al., 2021). Vegetation type strongly influenced the median  $g$  (Fig. 4b). We observe that evergreen broadleaf forests show the lowest value of 1 (Fig. 4b), whereas croplands have the highest value of 4.05 (Fig. 4b).

The spatial distributions of both  $L - VOD_{25km}$  (Fig. 4c) and the downscaled  $L - VOD_{5km}$  (Fig. 4e) in August 2021 exhibited  
255 consistent patterns. Notably high  $L - VOD_{5km}$  values were detected across tropical forests, with prominent examples being the Amazon and Congo basins; lower values were typical in short vegetation in tropical savannas and grasslands (e.g., Miombo woodlands in southern Africa); and very low values characterized arid and sparsely vegetated regions (e.g., northern Africa's Sahara and Australia's desert regions). In the median distributions across vegetation types,  $L - VOD_{25km}$  reached the highest value in EBF at 0.78 (Fig. 4d). Similarly,  $L - VOD_{5km}$  showed the highest value in EBF at 0.79 (Fig. 4f).

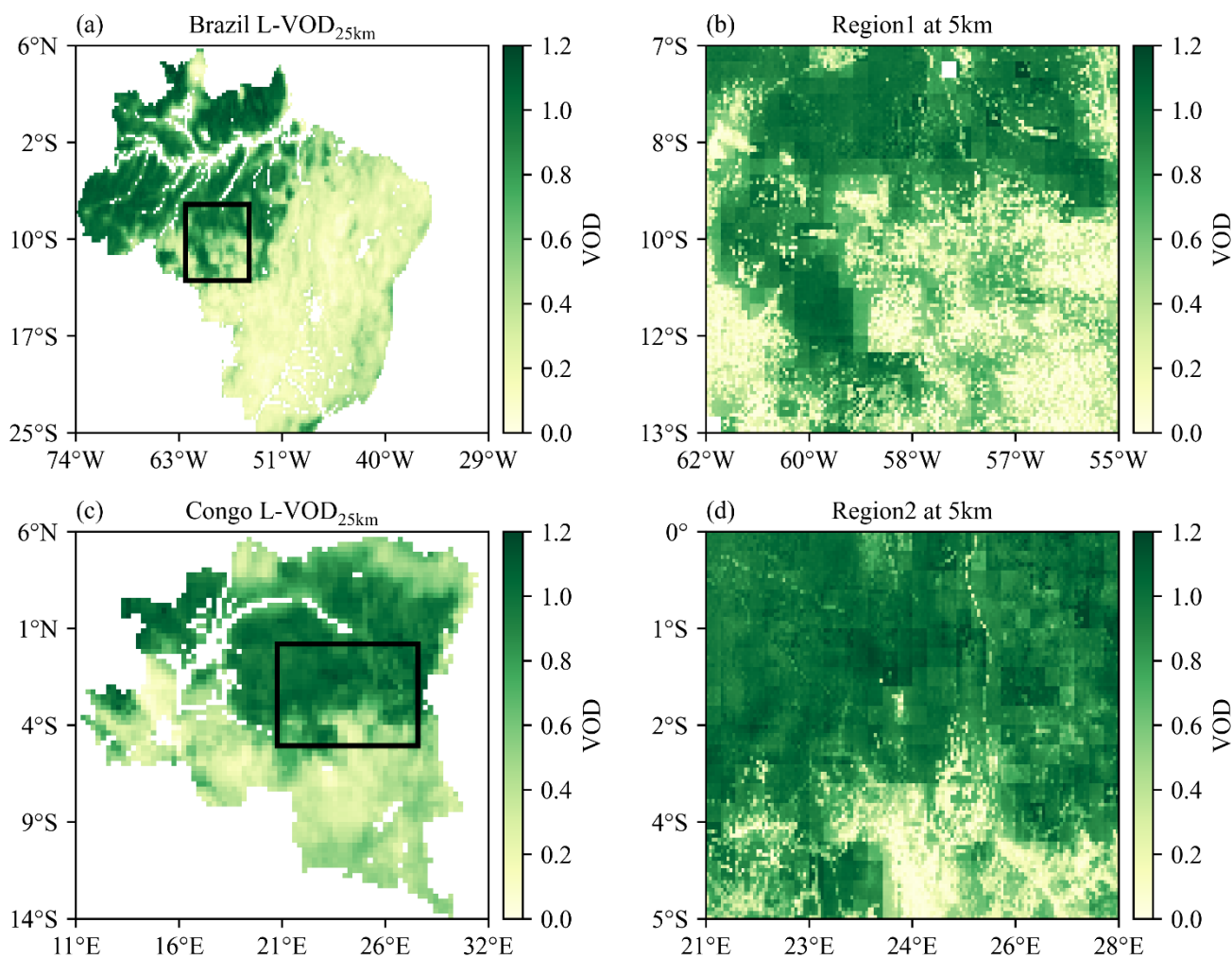


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**Figure 4. Spatial patterns and vegetation-class medians of the injection gains and L-VOD at 25-km and 5-km in August 2021. (a) Map of injection gains; (b) Median injection gains by vegetation class; (c) Map of  $L - VOD_{25km}$  (d) Median  $L - VOD_{25km}$  by vegetation class; (e) Map of  $L - VOD_{5km}$  (f) Median  $L - VOD_{5km}$  by vegetation class. Error bars denote the standard deviation of spatial variations.**

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Two representative sub-regions in Fig. 5(a, c) are provided as close-up views in Fig. 5b and Fig. 5d. In these areas, downscaled  $L - VOD_{5km}$  reveals more spatial details than the  $L - VOD_{25km}$ . For example, in Region1 (Fig. 5b), the 5 km L-VOD exhibits alternating high- and low-value patches, which likely reflect smaller-scale variations in vegetation water content or canopy structure. This indicates that the downscaling method effectively enhances the detailed spatial information, which is critical for capturing more localized vegetation dynamics.



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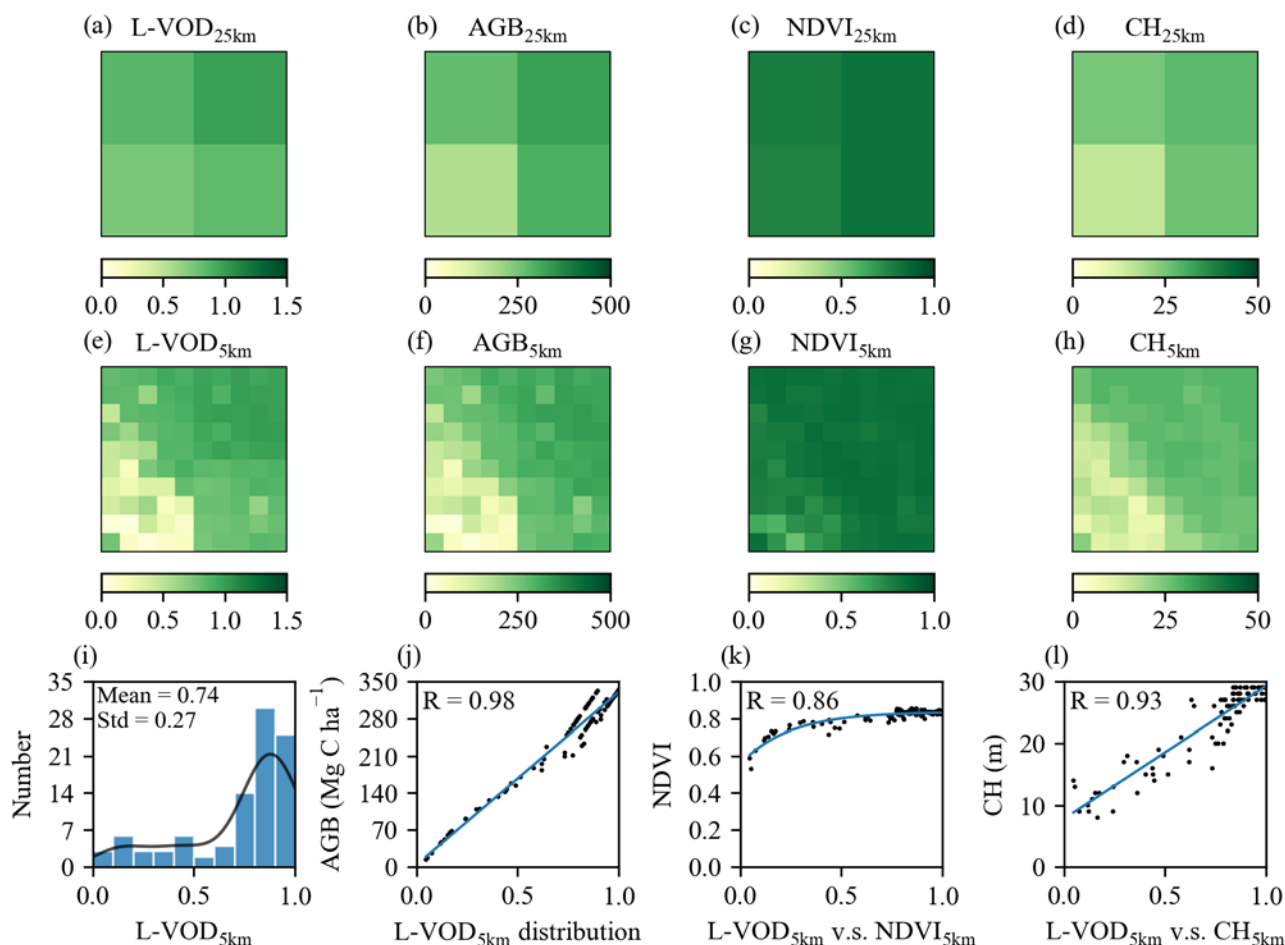
**Figure 5. Detailed spatial information of the downscaled L-VOD in August 2021. (a) 25-km resolution L-VOD of Brazil. (b) 5-km resolution L-VOD of Region 1 (the black rectangular area in (a)). (c) 25-km resolution L-VOD of the Congo. (d) 5-km resolution L-VOD of Region 2 (the black rectangular area in (c)).**

To further analyze the spatial details of  $L - VOD_{5km}$ , a small region of  $2 \times 2 L - VOD_{5km}$  pixels was selected (Fig. 6a).  
275  $L - VOD_{5km}$  shows the enhanced pixel-level heterogeneity within the selected region, particularly in the southwestern part (Fig. 6e). The pixel value range within the study area expanded substantially from 0.74-1.00 (Fig. 6a) for the  $L - VOD_{25km}$  to 0.05-1.01 for the  $L - VOD_{5km}$  (Fig. 6e). This indicates that the downscaling process effectively enhances spatial details and reveals finer variations that are not captured at the  $L - VOD_{25km}$ .

The distributions of  $L - VOD_{5km}$  (Fig. 6e) with AGB (Fig. 6f), NDVI (Fig. 6g), and CH (Fig. 6h) were found to be spatially  
280 similar within the selected  $2 \times 2$  pixels region. For all four variables, high values are concentrated in the southwestern part and low values in the north-western part.  $L - VOD_{5km}$  showed strong spatial correlations with AGB, NDVI, and CH, with



corresponding coefficients of 0.98 (Fig. 6j), 0.86 (Fig. 6k), and 0.93 (Fig. 6l), respectively, indicating the high similarity between  $L - VOD_{5km}$  and AGB.



285 **Figure 6. Detailed spatial information and relationships of L-VOD with AGB, NDVI and CH at 25-km and 5-km within four SMOS grid cells (corresponding to the locations of P1 in Fig. 1) in 2021. (a)  $L - VOD_{25km}$ ; (b)  $AGB_{25km}$ ; (c)  $NDVI_{25km}$ ; (d)  $CH_{25km}$ ; (e)  $L - VOD_{5km}$ ; (f)  $AGB_{5km}$ ; (g)  $NDVI_{5km}$ ; (h)  $CH_{5km}$ ; (i) Distribution of  $L - VOD_{5km}$ ; (j) Scatter of  $L - VOD_{5km}$  and  $AGB_{5km}$ ; (k) Scatter of  $L - VOD_{5km}$  and  $NDVI_{5km}$ , (l) Scatter of  $L - VOD_{5km}$  and  $CH_{5km}$ .**

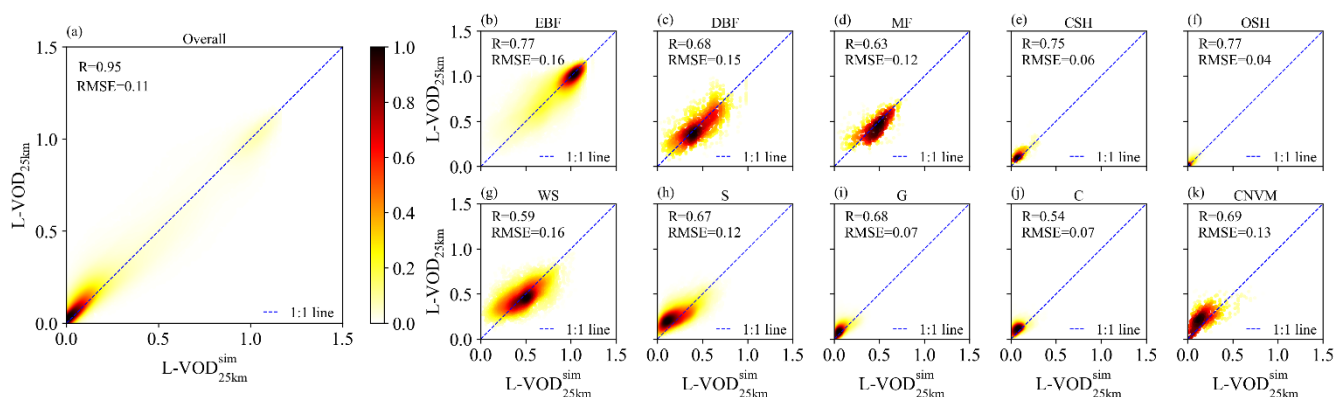
#### 4.2 Performance of the L-VOD downscaling method

290 The performance of the downscaling approach was assessed by calculating the  $R$  and  $RMSE$  between the  $L - VOD_{25km}^{sim}$  and the  $L - VOD_{25km}$  for the year 2019. The  $R$  between the  $L - VOD_{25km}^{sim}$  and  $L - VOD_{25km}$  reached 0.95 and  $RMSE$  reached 0.11 (Fig.7a). The findings highlight the capacity of our framework to conserve the coarse-resolution patterns of the original L-VOD product, while maintaining reliability and precision in evaluations involving spatial scaling.

The superior performance was observed in forests compared to all other vegetation types, particularly in evergreen broadleaf  
 295 forests (EBF), which exhibited the strongest correlation with  $R = 0.77$  and  $RMSE = 0.16$  (Fig. 7c). Conversely, croplands



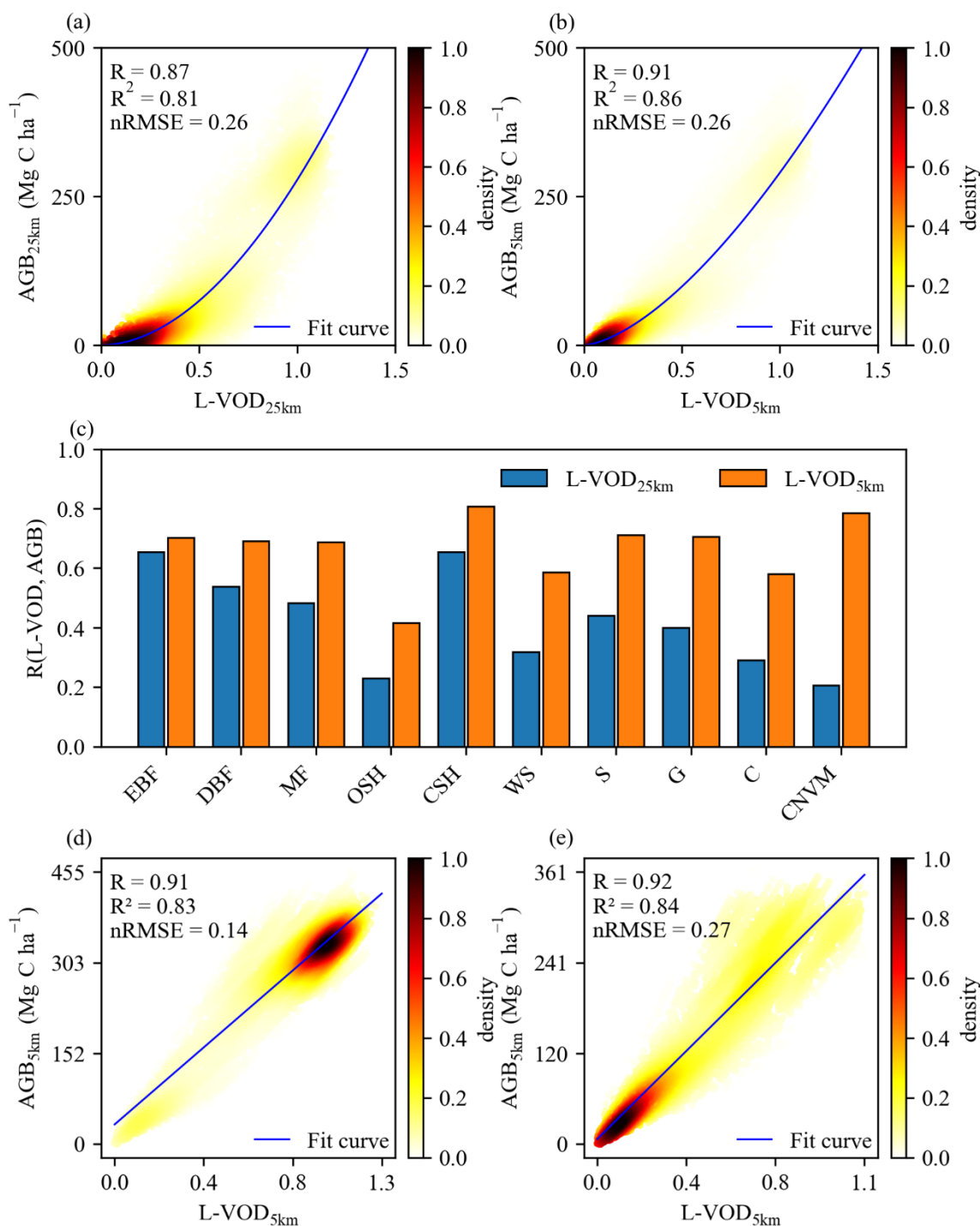
demonstrated the weakest correlation with  $R = 0.54$  and  $RMSE = 0.07$  (Fig. 7c). Notably, although the model achieves the same  $R$  values between  $L - VOD_{25km}^{sim}$  and  $L - VOD_{25km}$  for both EBF and OSH at  $R = 0.77$  (Fig. 7b, f), it yields a lower RMSE in OSH with  $RMSE = 0.04$  (Fig. 7f).



300 **Figure 7. Scatter density plots evaluating the proposed method overall and by vegetation class. (a) Overall performance between  $L - VOD_{25km}^{sim}$  and  $L - VOD_{25km}$ ; (b-k) Performance by IGBP vegetation classes—EBF, DBF, MF, CSH, OSH, WS, S, G, C, and CNVM (in order).**

### 4.3 The temporal and spatial assessment of downscaled L-VOD retrievals

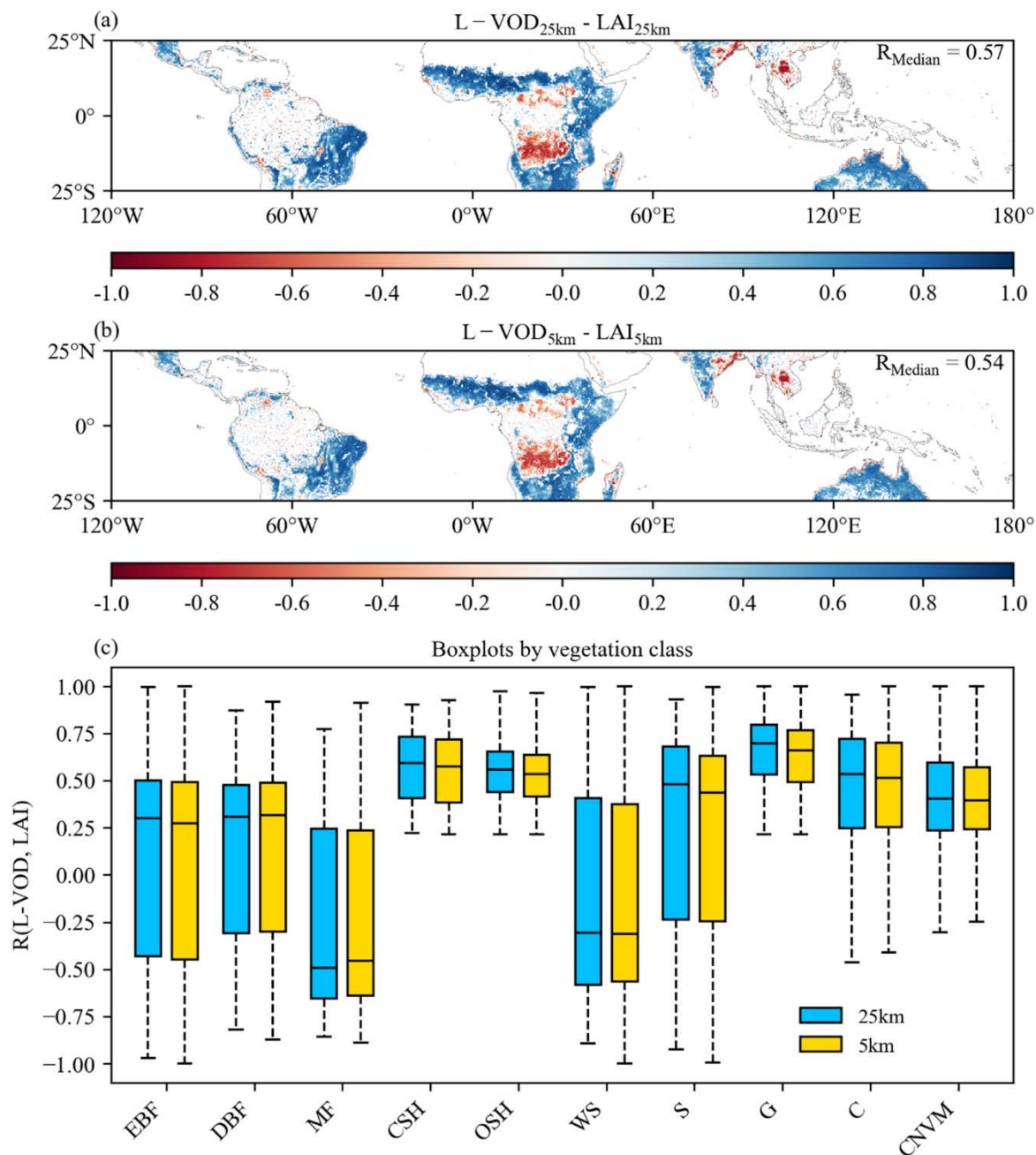
Further analysis revealed that the downscaled  $L - VOD_{5km}$  achieved an even higher spatial  $R$  of 0.91 with  $AGB_{5km}$  (Fig. 8b),  
 305 relative to 0.87 (Fig. 8a) in 2019. The improvement was also reflected by the coefficient of determination, with  $R^2$  increasing  
 from 0.81 to 0.86. This pattern was consistently observed across all other study years (Fig. A1). Meanwhile, the median of the  
 pixel-wise temporal correlation  $L - VOD_{5km}$  with  $AGB_{5km}$  reached  $R = 0.9$  (Fig. A2). This indicates that  $L - VOD_{5km}$   
 effectively integrates fine-scale spatial structural information, while also capturing the temporal dynamics of vegetation growth.  
 Furthermore, a universal improvement in  $R$  is observed across all land cover types. For instance, three major land cover  
 310 categories—EBF, G, and CSH—exhibited  $R$  value improvements: EBF increased from 0.65 to 0.70 (Fig. 8c), G from 0.40 to  
 0.70 (Fig. 8c), and CSH from 0.651 to 0.81 (Fig. 8c). EBF shows the smallest improvement, while CNVM shows the largest.  
 We analyzed the spatial correlation between  $L - VOD_{5km}$  and  $AGB_{5km}$  in two representative regions: the Democratic Republic  
 of the Congo (DRC, Fig. 8d) and Rondônia (Fig. 8e). In the region of DRC,  $L - VOD_{5km}$  achieved a high  $R$  of 0.91 with  
 $AGB_{5km}$ , coupled with a low normalized root mean square error ( $nRMSE = 0.135$ ) and a high  $R^2$  (0.83). In the Rondonia,  
 315  $L - VOD_{5km}$  still maintained high correlation ( $R = 0.92$ ,  $R^2 = 0.84$ ) with  $AGB_{5km}$ , though  $nRMSE$  (0.27) was slightly elevated.  
 These results demonstrate that  $L - VOD_{5km}$  is capable of effectively capturing the subtle AGB gradients within local areas.



320 **Figure 8. Spatial relationships between annual average L-VOD and AGB at 25-km and 5-km resolutions in 2019. (a) Density scatterplot between  $L-VOD_{25km}$  and  $AGB_{25km}$ ; (b) Density scatterplot between  $L-VOD_{5km}$  and  $AGB_{5km}$ ; (c) Comparison median value of  $R$  of L-VOD and AGB for each vegetation type; (d) Density scatterplot between  $L-VOD_{5km}$  and  $AGB_{5km}$  in the DRC; (e) Density scatterplot between  $L-VOD_{5km}$  and  $AGB_{5km}$  in Rondonia.**



The temporal correlations of L-VOD with LAI remained at a comparable level before and after downscaling (Fig. 9a, b), suggesting that our downscaling approach maintains the temporal dynamics of the original L-VOD. This was indicated by that the median  $R$  between the  $L - VOD_{25km}$  and LAI was 0.57 (Fig. 9a), consistent with the 0.54 (Fig. 9b) correlation between  $L - VOD_{5km}$  and LAI. This was further confirmed by the similar values of temporal correlation between LAI,  $L - VOD_{5km}$  and  $L - VOD_{25km}$  over different vegetation types (Fig. 9c). For example, in grasslands,  $R$  was 0.7 for  $L - VOD_{25km}$  and 0.66 for  $L - VOD_{5km}$  (Fig. 9c). A strong similarity was observed between the spatial patterns of L-VOD—NDVI temporal correlation (Fig. A3) and that of LAI. In general,  $L - VOD_{5km}$  correlated better with NDVI ( $R = 0.57$ ) than with LAI ( $R = 0.54$ ).



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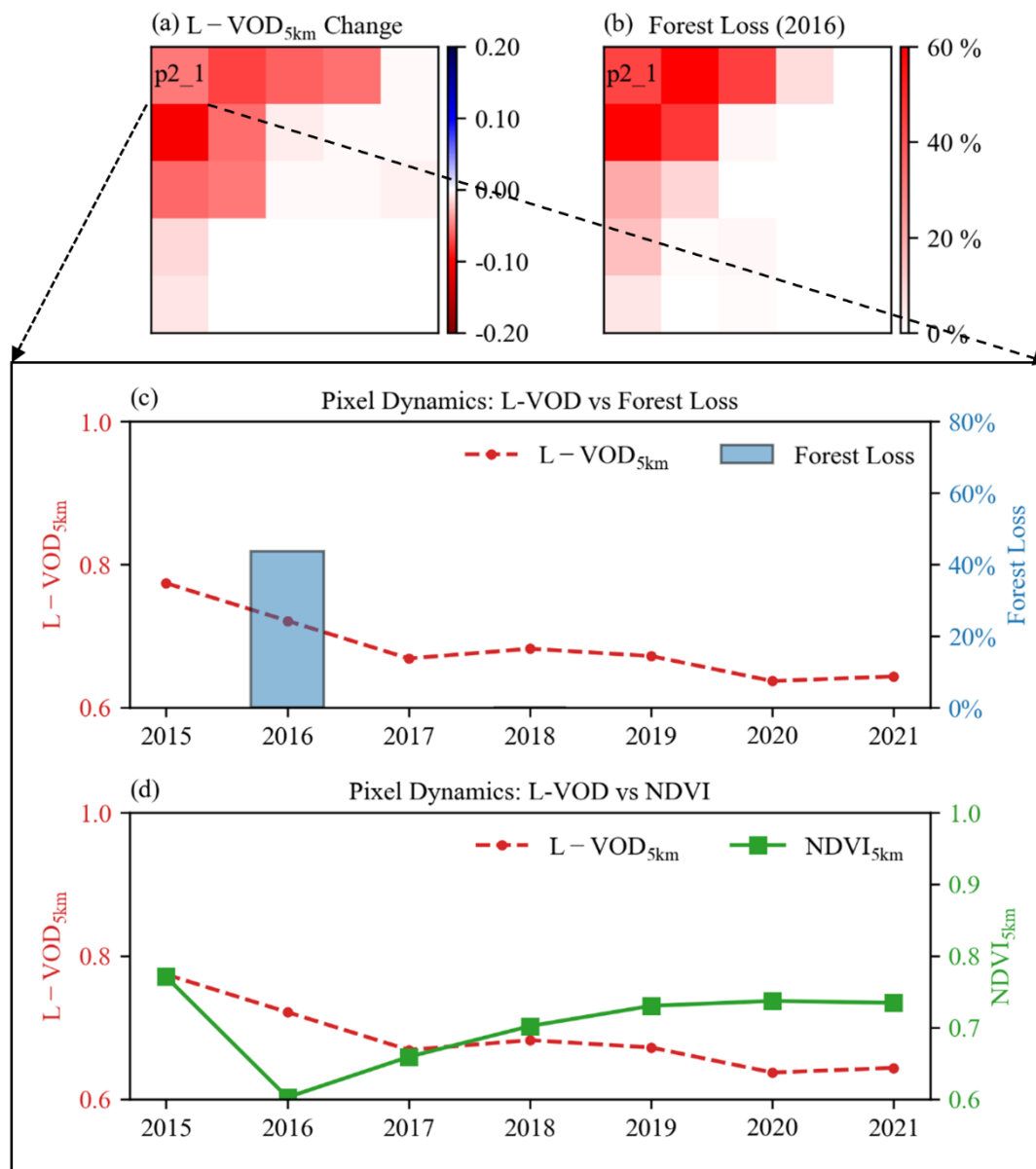
**Figure 9.** Spatially distributed temporal correlation coefficients ( $R$ ) between monthly L-VOD and LAI from 2015 to 2021 at 25-km resolution (a) and at 5-km resolution (b). (a) and (b) blank areas represent pixels that were filtered out by scene flags, as well as desert regions and statistically non-significant correlations ( $p > 0.05$ ).

We evaluate the  $L - VOD_{5km}$  product with forest loss data, given that L-VOD has been shown to capture forest loss at 25-km resolution (Qin et al., 2021). The spatial distribution of forest loss area (Fig. A4) shows that the Brazilian Amazon contains

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areas of both high loss intensity and clear spatial clustering. We selected one grid cell for in-depth analysis in the Brazilian Amazon located at P2 (marked in Fig. 1), which experienced a 26% forest loss mainly concentrated in the north-western area during 2016. Within 25 km scale,  $L - VOD_{5km}$  pixels exhibited a pronounced decline generally in the northwestern part of the cell (Fig. 10a). Correspondingly, the percentage (45%–64%) of forest loss area within the grid was also concentrated in the north-western region (Fig. 10b), indicating a strong spatial consistency between the changes in  $L - VOD_{5km}$  and forest loss. In terms of temporal dynamics,  $L - VOD_{5km}$  effectively capture changes in forest loss. A 5km × 5km forest sample was selected at the site indicated by P2\_1 in Fig. 10a, b. The L-VOD of this sample exhibited its largest decline between 2015 and 2016, decreasing from 0.77 to 0.72, a trend that continued into 2017 (Fig. 10c), while forest loss during the same period increased from 0% to 44% (Fig. 10c). This may be attributed to carbon losses captured by L-VOD across the Brazilian Amazon being likely influenced by additional drivers beyond deforestation in a given year (Li et al., 2022b). For instance, forest degradation can lead to a decline in L-VOD, as verified by Qin et al. (2021). Additionally, we compared  $L - VOD_{5km}$  temporal dynamics with  $NDVI_{5km}$  in the same sample (Fig. 10d). In 2015,  $L - VOD_{5km}$  and NDVI are the same value 0.77, after the 2016 forest disturbance,  $NDVI_{5km}$  dropped sharply from 0.77 to 0.60 but underwent rapid recovery in subsequent years (approaching 0.73 by 2021). In contrast,  $L - VOD_{5km}$  also declined following the 2016 disturbance, but failed to recover rapidly thereafter. The lag in  $L - VOD_{5km}$  was ascribed to the relatively slower recovery of both non-photosynthetic (woody) and photosynthetic (foliar) canopy biomass, in contrast to the more rapid canopy greenness response detected by  $NDVI_{5km}$  (Jones et al., 2013). This demonstrates that the downscaled  $L - VOD_{5km}$  not only captures vegetation disturbance but also reveals the lagged recovery of canopy biomass, effectively complementing the information provided by optical vegetation indices.



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Figure 10. Spatiotemporal consistency analysis between downscaled L-VOD and forest loss within a 25-km grid cell in the Brazilian Amazon (corresponding to the locations of P2 in Fig. 1). (a) Change in L-VOD from 2015 to 2016 at 5-km resolution; (b) Forest loss percentages in 2016 at 5-km resolution; (c) Pixel-level (corresponding to the locations of P2\_1 Fig. 9a, b) temporal dynamics of L-VOD (left axis) and forest loss area (right axis) from 2015 to 2021 at 5-km resolution; (d) temporal dynamics of L-VOD (left axis) and NDVI area (right axis) from 2015 to 2021 at 5-km resolution.

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## 5 Discussion

### 5.1 Overview of the strengths of the downscaling framework

This study introduces a framework for 5-km L-VOD estimation. Unlike statistical regression methods, our method sharpens  $L - VOD_{5km}$  by injecting spatial structural features (e.g., texture, edges) from  $AGB_{100}$  data into the  $L - VOD_{25km}$  baseline, rather than relying on auxiliary data absolute values. A direct jump to 1-km resolution ( $25 \times 25$  expansion) would generate unconstrained spatial details that lack support from  $L - VOD_{25km}$ , causing severe spatial distortions and loss of physical interpretability. Furthermore, our method is intended to integrate the respective advantages of the SMOC L-VOD products and CCI-AGB datasets: the temporal resolution of  $L - VOD_{25km}$  and the spatial resolution of  $AGB_{100m}$ , and its potential benefits relative to conventional methods are explored in the following sections.

This method preserves the original temporal resolution of the VOD because it uses the VOD as a baseline for structural-feature injection. In comparison, the traditional downscaling approach is unable to eliminate the impacts of temporal discrepancies, since other auxiliary satellite datasets usually possess inconsistent observation times. (Hongtao et al., 2019). Furthermore, it does not require other ancillary satellite data that are susceptible to cloud contamination (such as optical/thermal infrared data), so the error in the  $L - VOD_{5km}$  estimation process can be decreased.

The downscaling framework is to estimate higher-resolution VOD images by injecting the structural information from higher-resolution auxiliary datasets, instead of establishing explicit or implicit functions to convert pixel values from other satellite products (Bousquet et al., 2019; Mohite et al., 2022; Yan et al., 2025). Furthermore, our framework does not rely on a retrieval process, thereby avoiding the uncertainties within the radiative transfer models (RTMs) (Vergopolan et al., 2020).

To evaluate the performance of this approach, we compared it with the Gram–Schmidt adaptive (GSA) (Aiazzi et al., 2007), which is based on a linear transformation (Method B1). The proposed downscaling model achieved  $R$  of 0.95 Fig. 6a, whereas the GSA method showed a lower  $R$  at 0.86 (Fig. A5a). Fig. A5b, c presents a comparison of spatial details between the original L-VOD and the downscaled L-VOD products derived from two approaches across the study areas. It is evident that the downscaled L-VOD obtained using the method proposed in this study shows higher consistency in textural characteristics with the original L-VOD, as compared with the GSA method.

### 5.2 Uncertainties and perspectives

Despite the promising results, the framework still has some certain limitations. In the following part, we analyze the influences of these uncertainties and outline possible pathways for future improvement.

The injection gains  $g$  (see Eq. (2)), shown in Fig. 4a, a key parameter in our downscaling process, had uncertainties, because it is derived from the  $L - VOD_{5km\_NNR}$  and  $AGB_{5km\_Low\_Pass}$ . Firstly, the  $L - VOD_{5km\_NNR}$  is generated from the original L-VOD products using the NNR method. While this method effectively preserves the original values, it may cause the loss of textural details (Parker et al. 1983). Additionally, it should be noted that these L-VOD already contain inherent errors introduced by satellite sensor imaging and the retrieval algorithms (Hu et al., 2024). Secondly, a fundamental assumption in



our methodology may introduce another potential source of error: the CCI-AGB data, which provide the static higher-resolution spatial structure, are assumed to be constant on a monthly basis within each year, as AGB represents a relatively stable structural component of vegetation, whereas the seasonal dynamics are primarily captured by VOD. This simplification may fail to capture intra-annual biomass dynamics, potentially leading to discrepancies when fused with monthly L-VOD observations. Furthermore, as the CCI-AGB dataset only characterizes forest biomass, its application to non-forest regions during the downscaling process may introduce additional errors in the extracted spatial features. Accordingly, future research efforts will prioritize the incorporation of high-resolution vegetation structural datasets with comprehensive coverage across diverse vegetation types, which is expected to mitigate such vegetation-type-related biases and improve the robustness of the downscaling framework across heterogeneous landscapes.

Direct validation of  $L - VOD_{5km}$  is limited by the absence of consistent *in situ* observations or reliable model outputs (Li et al., 2020). We evaluated the temporal performance of  $L - VOD_{5km}$  by employing two optical vegetation-related indices, namely NDVI and LAI, following the method proposed by Li et al. (2022a). Nevertheless, the variability discrepancy between  $L - VOD_{5km}$  and VIs may stem from the fact that L-VOD represents vegetation water content (VWC), which conveys distinct information from optical VIs (NDVI shows greater sensitivity to leaf chlorophyll content, while LAI characterizes the overall density of vegetation). This could induce temporal lags and dynamic variations between their time series (Zhou et al., 2022). Acknowledging these limitations, we still believe that the chosen indicators are appropriate for evaluating  $L - VOD_{5km}$  in present study, given that multiple studies have confirmed that strong spatiotemporal consistency between optical VIs and VOD, most notably in regions dominated by low-stature vegetation (Li et al., 2020; Li et al., 2022a; Li et al., 2021). Uncertainties associated with optical VIs may partially account for the weak correlation between these indices and  $L - VOD_{5km}$  across portions of the study area (Forkel et al., 2013). Cloud interference and aerosol particles, for instance, often degrade data quality and introduce gaps in daily time series, yielding VIs that do not reliably reflect actual vegetation variability.

Recent technological advances are providing new opportunities to validate VOD more rigorously. First, field configurations using global navigation satellite systems (GNSS) can measure VOD using cost-effective sensors that receive microwave signals from navigation satellites. Second, land surface models are now including methods to model vegetation water content and thus estimate VOD globally at the same spatial scales as satellite measurements (Feldman, 2024). Therefore, the further development of these GNSS and model methods are needed such that satellite VOD can be more comprehensively validated.

## 6 Conclusion

In this study, SMOS L-VOD data were spatially enhanced from original 25-km resolution to 5-km over tropical regions using a downscaling L-VOD framework, which fuses higher-resolution structural information from AGB to enhance spatial details of the original SMOS L-VOD data.

In the absence of ground-based L-VOD observations, three complementary assessment approaches were implemented to examine the reliability of the 5-km products across tropical regions, as follows: (1) spatial evaluation using AGB ( $R=0.91$ ,



425  $R^2 = 0.86$ ), (2) pixel-level temporal consistency assessment using LAI and NDVI time series, and (3) disturbance response analysis based on forest loss data. Quantitative assessments demonstrate that the downscaled datasets effectively preserve the broad-scale spatial patterns of the original input ( $R = 0.95$  and  $RMSE = 0.11$ ) while successfully fusing finer-scale structural information. The product also shows improved capacity in capturing localized variations in vegetation dynamics and demonstrates sensitivity to forest disturbances.

430 **Appendix A**

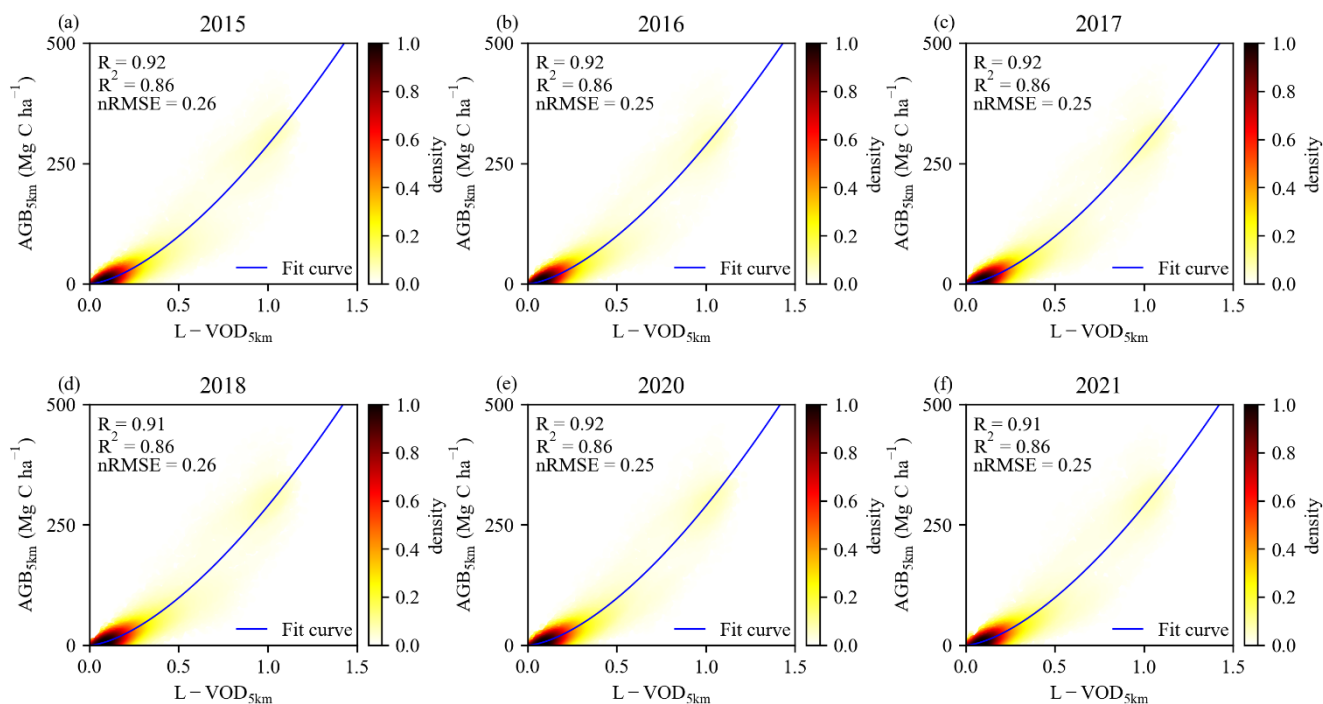
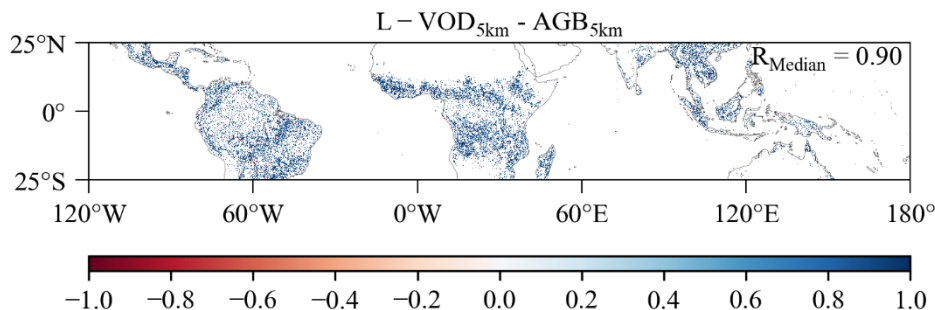
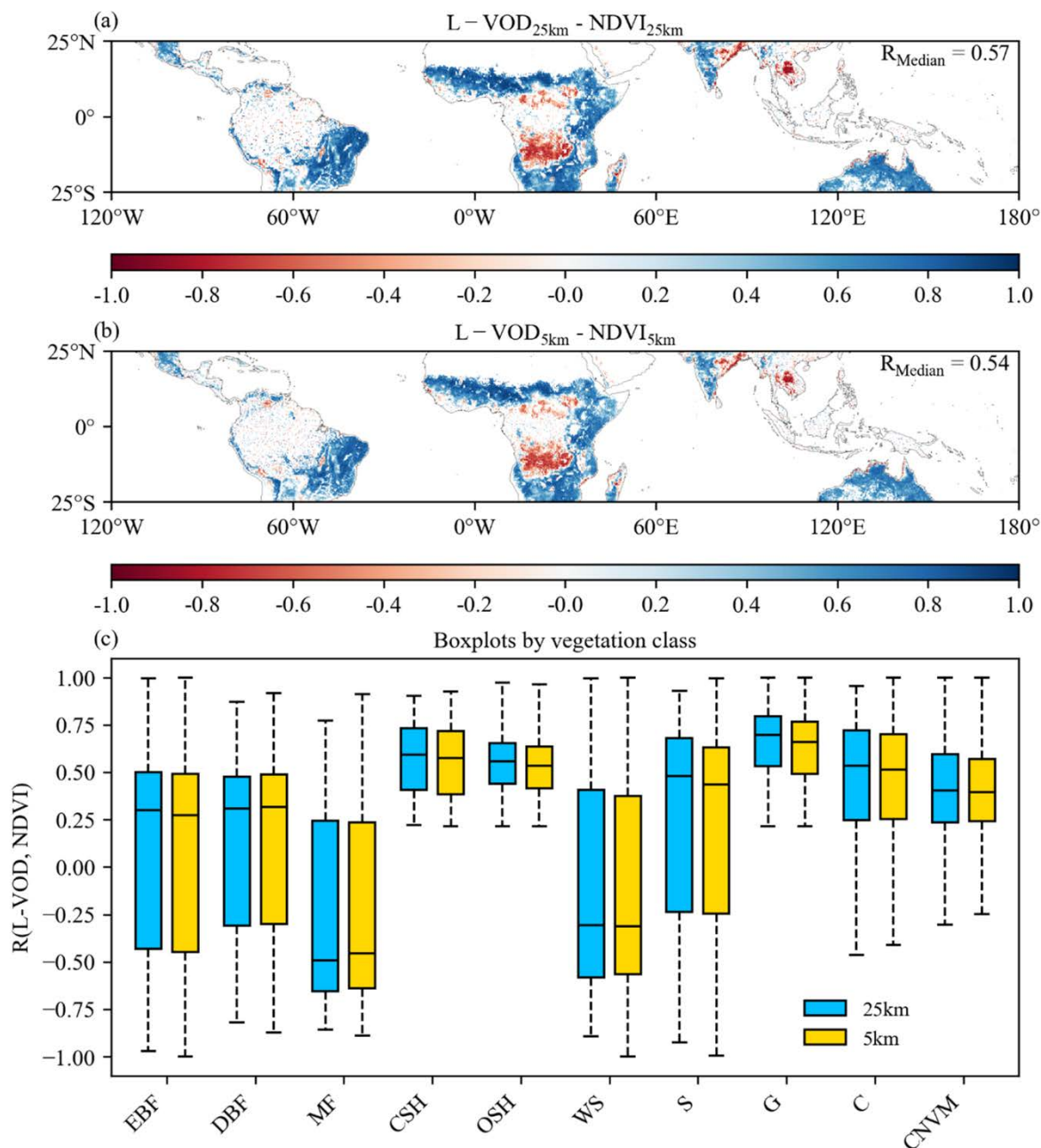


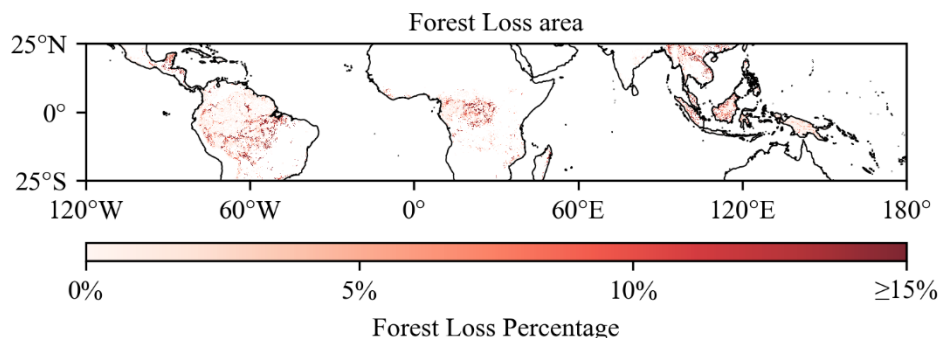
Figure A1. Correlations between yearly  $L - VOD_{5km}$  and  $AGB_{5km}$  from 2015 to 2018, and in 2020 and 2021.



435 Figure A2. Spatial distribution of temporal correlation coefficients between  $L - VOD_{5km}$  and  $AGB_{5km}$  over the 2015–2021 period. White areas indicate pixels with either missing valid data or statistically non-significant relationships ( $p > 0.05$ ).

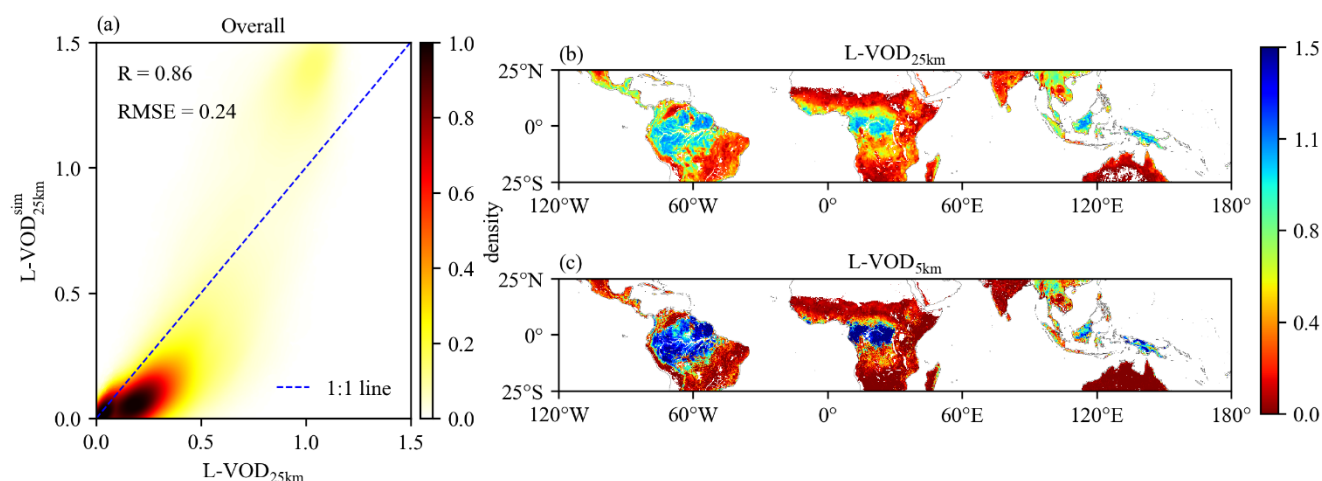


**Figure A3.** Pixel-wise temporal correlation ( $R$ ) between monthly L-VOD and NDVI from 2015 to 2021 at 25-km resolution (a) and 5-km resolution (b); (c) Boxplots of correlations by IGBP vegetation class at 25-km and 5-km resolutions. (a) and (b) blank areas represent pixels that were filtered out by scene flags, as well as desert regions and statistically non-significant correlations ( $p > 0.05$ ).



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**Figure A4.** Maps of 5-km resolution forest loss area fraction for 2015–2021.



**Figure A5.** Results of GSA-based downscaling of L-VOD. (a) Density scatterplot comparing  $L - VOD_{25km}^{sim}$  and  $L - VOD_{25km}$ ; (b) Spatial pattern of the  $L - VOD_{25km}$ ; (c) Spatial pattern of the  $L - VOD_{5km}$ .

## 445 Appendix B

### Method B1. Gram-Schmidt Adaptive (GSA) Downscaling Method

The Gram-Schmidt Adaptive (GSA) method, a widely recognized pan-sharpening technique (Aiazzi et al., 2007), was employed as a benchmark to contextualize the performance of our developed downscaling method. In this study, we adapted the core principle of GSA—component substitution guided by multivariate regression—to the specific task of downscaling L-  
 450 VOD data. The downscaling process involved three steps:

- (1) Regression at Low Resolution: A multivariate regression model  $L - VOD_{25km} = f(AGB_{25km})$  was established to learn the adaptive linear relationship between the two variables at the native 25-km scale.
- (2) Spatial Component Synthesis: The derived regression coefficients were applied to the native  $AGB_{5km}$  to synthesize a higher-resolution spatial component  $I_{5km}$ , which inherited the fine spatial structure of AGB.



455 (3) Fusion: The spatial details from  $I_{5km}$  were injected into an up-sampled version of the original L-VOD data via a component substitution scheme to generate the final  $L - VOD_{5km}$ .

### Data availability

The pan-tropical 5-km monthly L-band vegetation optical depth dataset (2015–2021) is available at <https://doi.org/10.11888/RemoteSen.tpd.303391> (Shi and Fan, 2026).

### 460 Author contributions

JS and LF designed the experiments, and JS carried them out. LF, JL, LG, QY and SL provided related suggestions. LF, RF, PC, XL and SL reviewed and edited the text. JS prepared the manuscript with contributions from all co-authors.

### Competing interests

The authors declare that none of the authors has any conflict of interest.

### 465 Financial support

This work was financially supported by the National Natural Science Foundation of China (42322103), the Chongqing Outstanding Youth Science Foundation (CSTB2024NSCQ-JQX0010).

### Review statement

470 The review statement will be added by Copernicus Publications listing the handling editor as well as all contributing referees according to their status anonymous or identified.

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