Machine-learning-based estimates of global natural vegetated wetland methane emissions (2000-2025)

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Abstract. Wetlands are the largest natural source of atmospheric methane (CH_4), yet comprehensive global budgets are typically delayed by years, preventing a timely understanding of CH_4 sources, sinks, and trends. To reduce this delay, we present a model emulator-driven framework and accompanying workflow that enable timely, continuous emission updates using a machine-learning emulator to reconstruct spatially explicit monthly emission fields at 1°x1° resolution. We apply this framework to a global dataset of natural vegetated wetland CH_4 emissions to extend the most recent Global Methane Budget (GMB; Saunois et al., 2025) record that covers the 2000-2020 emissions through 2025. In the test data (~30% of the total dataset), the emulator achieved a global R^2 of 0.65 ± 0.003 (mean ± 95% CI, hereafter) and an RMSE of 5.49 ± 0.12 $\times 10^{-3}$ Tg CH_4/year. The emulator is trained on 35 GMB model estimates, including 22 process-based models and 13 atmospheric inversions, paired with 10 ensemble realizations of 11 gridded climate predictor variables from atmospheric reanalyses. Our results show

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that the global mean predicted wetland CH_4 emissions for 2021-2025 (157.8 ± 2.4 Tg CH_4/year) are not significantly higher (~0.05 Tg CH_4/year) than the 2000-2020 baseline. However, this stability masks a significant hemispheric redistribution of emissions. We detect an increase in Northern Hemisphere (NH) emissions in 2021-2025, with mid- and high-latitudes increasing by 0.76 ± 0.07 and 0.35 ± 0.03 Tg CH_4/year, respectively, while the tropics and Southern Hemisphere (SH) extratropics show offsetting negative trends (-0.95 ± 0.19 and -0.11 ± 0.02 Tg CH_4/year, respectively). The predicted emissions are able to capture the low emissions in 2023 in South America linked to El Niño-related drought, as reported by recent studies (Ciais et al., 2026; Quinn et al., 2025). Furthermore, we identify a distinct seasonal amplification of global emission trends that peaks in late boreal summer. This new modeled dataset and operational framework bridge the gap between the latest updated budgets and low-latency monitoring, providing a scalable capacity to frequently update global emission estimates and critical early warnings of regional wetland feedback loops. The data are publicly available at https://doi.org/10.5281/zenodo.18870108 (Li et al., 2026a).

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1 Introduction

Wetlands are the largest natural source of methane (CH_4), a potent greenhouse gas, yet they remain one of the most uncertain components of the global CH_4 budget (Saunois et al., 2025; Zhu et al., 2025). They account for 25-30% of global CH_4 emissions and strongly influence the global carbon cycle (Jackson et al., 2024; Li et al., 2026b; Peng et al., 2022; Saunois et al., 2025; Zhang et al., 2023). Quantifying how wetland CH_4 emissions respond to climate variability is essential for interpreting interannual variability in the global CH_4 budget and for understanding recent atmospheric CH_4 changes (Ciais et al., 2026; Gedney et al., 2019a; Parker et al., 2018; Stavert et al., 2022). Community efforts such as the Global Methane Budget (GMB) synthesize process-based "bottom-up" (BU) models and atmospheric "top-down" (TD) inversions to estimate all CH_4 emission sources and sinks (Saunois et al., 2020, 2025). However, the substantial effort required to assemble these products into a reliable ensemble typically leads to lags of multiple years. For instance, the most recent GMB published in 2025 (Saunois et al., 2025) covered the period 2000-2020. Although low-latency estimates exist for some individual models (Quinn et al., 2025), accelerating the assembly of a trustworthy multi-model synthesis would better support timely emission monitoring and evidence-based assessment of mitigation progress.

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Wetland CH_4 emissions combine biogeochemical and hydrometeorological controls that regulate CH_4 production with oxidation processes within sediments and the water column. Temperature exerts a strong influence on microbial methanogenesis and ecosystem respiration (Knox et al. 2019; Bansal et al., 2023; Li et al., 2025; Yvon-Durocher et al., 2014), whereas soil thermal and free-thaw conditions can modulate growing-season length and cold-season CH_4 dynamics (Hyvärinen et al., 2025; Zona et al., 2016). Hydrology, often summarized by water table depth, soil moisture content, and inundation extent, controls oxygen availability and redox conditions and influences wetland CH_4 production and oxidation (Cui et al., 2024; He et al., 2025; Knox et al., 2021). Precipitation, lateral water transport, and

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surface energy fluxes jointly constrain water balance through inputs and evapotranspiration (Aalto et al., 2025; Helbig et al., 2020; Tyystjärvi et al., 2024). Furthermore, radiation and vegetation state provide proxies of substrate supply and plant-mediated transport pathways that connect below-ground CH_4 production to the atmosphere (Helfter et al., 2022; McNicol et al., 2023).

Currently, global wetland CH_4 emissions are primarily estimated using two approaches: BU and TD models. Process-based BU models explicitly simulate the complex biogeochemical mechanisms driven by environmental conditions (Zhang et al., 2025). TD atmospheric inversions estimate surface fluxes by optimising prior emission inventories to match observed atmospheric CH_4 concentrations (Patra et al., 2018). Recent studies have increasingly turned to data-driven machine-learning models. For instance, McNicol et al. (2023) upscaled eddy covariance CH_4 fluxes globally using random forest algorithms; Bernard et al. (2025) introduced a satellite observation-based model to simulate temporal emission variability. These machine-learning approaches offer some advantages, including rapid operational speeds and lower computational costs. However, they are constrained by the sparsity of training data, particularly in the tropics and high latitudes.

Building on these well-established controls and the advancements in data-driven modeling, we provide a low-latency continuation of the GMB gridded wetland CH_4 emissions by extending monthly emissions from 2000-2020 through 2025 at 1°x1° resolution. Our approach is supported by the fact that many key climate, hydrometeorological and vegetation drivers are routinely updated as global gridded datasets, enabling a practical emulator-based pathway for timely wetland CH_4 emission estimates. In this study, wetlands are defined following the GMB natural vegetated wetland category, excluding lakes, rivers, reservoirs, coastal waters and managed sources (Saunois et al., 2025). We emulate each BU and TD ensemble member from GMB using machine-learning models trained on the 2000-2020 GMB wetland emission fields and climate predictor variables (e.g., soil temperature and precipitation) from ERA5 reanalysis to yield spatially explicit monthly emissions for 2000-2025 with uncertainty derived from both GMB ensemble spread (35 runs) and ERA5 monthly reanalysis (10 ensemble members). Using the resulting 2000-2025 wetland CH_4 emission data, we then quantify recent emission changes in 2021-2025 relative to 2000-2020, characterize interannual variability in annual regional and latitudinal emission time series, and diagnose long-term growth rates and their seasonality.

Our proposed framework and dataset provide a lower-latency, scalable tool for spatially explicit wetland CH_4 emission estimates. It provides support for early-warning diagnostics, attribution of recent atmospheric CH_4 anomalies, and provision of timely priors for atmospheric inversions.

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2 Methods

2.1 Input datasets

2.1.1 Wetland CH_4 flux data

We used monthly net CH_4 flux estimates for natural vegetated wetlands from the most recent Global Methane Budget (GMB) synthesis for 2000-2020, including both process-based BU models and atmospheric TD inversions (Saunois et al., 2025; Zhang et al., 2025). In this study, the natural vegetated wetland definition excludes lakes, rivers, reservoirs, coastal waters, and managed sources, following the underlying GMB definition (Saunois et al., 2025). Detailed descriptions of each model and inversion framework are provided in the cited GMB publications; here we summarize the key elements relevant to this study.

The BU wetland biogeochemical models (22 estimates, Table S1) included in the analysis are from 11 models: CLASSIC (Arora et al., 2018; Melton and Arora, 2016), ELM-ECA (Riley et al., 2011), ISAM (Shu et al., 2020; Xu et al., 2021), JSBACH (Kleinen et al., 2020, 2021, 2023), JULES (Gedney et al., 2019b), LPJ-MPI (Kleinen et al., 2012), LPJ-WSL (Zhang et al., 2016), LPX-Bern (Spahni et al., 2011; Stocker et al., 2014), ORCHIDEE (Ringeval et al., 2011), SDGVM (Beerling and Woodward, 2001; Hopcroft et al., 2011, 2020), and VISIT (Ito and Inatomi, 2012). Each model was run under two global climate forcings (CRU and GSWP3-W5E5). All BU simulations are prognostic, i.e. each model computes wetland extent internally rather than relying on a shared wetland extent dataset. Differences in simulated wetland extent therefore contribute to inter-model spread in CH_4 emissions and may explain a substantial fraction of discrepancies across regions.

The TD inversion products (13 estimates, Table S2) were constrained by surface or satellite observations over the 2000-2020 period and they are from seven systems: CarbonTracker Europe-CH_4 (Tsuruta et al., 2017), LMDz-CIF (Thanwerdas et al., 2022), LMDz-PYVAR (Zheng et al., 2018a, b, 2019), MIROC4-ACTM (Chandra et al., 2021; Patra et al., 2018), NISMON-CH_4 (Niwa et al., 2022, 2025), and NIES-TM-FLEXPART (NFVAR) (Maksyutov et al., 2021; Wang et al., 2019). Two alternative anthropogenic prior inventories, EDGAR v6 (Crippa et al., 2021) and GAINS (Höglund-Isaksson et al., 2020), were used to represent anthropogenic emissions. Wetland and inland freshwater emissions were prescribed as separate prior fluxes in the inversion framework. Wetland priors were derived from the ensemble mean of dynamic process-based BU wetland models used in the GMB estimates, whereas inland freshwater emissions were based on independent BU estimates. We used posterior flux estimates throughout this study.

All BU and TD outputs provide monthly wetland CH_4 flux as emission per grid cell area, typically at 1°x1° resolution in the regridded GMB output products. Coarser products (CLASSIC models at 1.85° and JSBACH models at 2.8°) were remapped to a 1°x1° grid for spatial consistency across datasets. In total, we included 22 BU estimates and 13 TD estimates (35 estimates overall). Tables S1 and S2 summarize the BU and TD estimates.

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2.1.2 Predictor variables from ERA5 reanalysis

We assembled 10 ensemble members of global ERA5 monthly averaged data at 0.5° x 0.5° spatial resolution (Hersbach et al., 2023) for 11 climate predictor variables: t2m (near-surface air temperature), tp (total precipitation), ssrd (downward shortwave radiation at surface), slhf (surface latent heat flux), sshf (surface sensible heat flux), swvl1 (volumetric soil water in layers 1, 0-7 cm below surface), swvl2 (volumetric soil water in layers 2, 7-28 cm below surface), stl1 (soil temperature in layer 1, 0-7 cm below surface), stl2 (soil temperature in layer 2, 7-28 cm below surface), lai_hv (high-vegetation leaf area index), and lai_lv (low-vegetation leaf area index). LAI (leaf area index) from ERA5 is prescribed in the ECMWF land-surface model as a seasonally varying, monthly climatology derived from satellite LAI (Roberts et al., 2018). We chose ERA5 data in part because it is operationally updated regularly, which enables continuous extension of predictor fields and low-latency updating of emission reconstructions as new months become available. ERA5 data has also undergone extensive evaluation and is widely used due to its high spatiotemporal consistency and generally strong performance (Hersbach et al., 2020), but notably regional biases remain, for example, a weaker performance in parts of Asia/Africa for hydrology (Gebrechorkos et al., 2024) and temperature in the Arctic (Tian et al., 2024). The 10 ensemble members correspond to the ERA5 Ensemble of Data Assimilations, which provides multiple physically consistent realizations of the reanalysis by perturbing the data assimilation system to sample uncertainty in atmospheric and land-surface states. The ensemble spread can be interpreted as an indicator of reanalysis uncertainty.

The 11 variables included in this analysis represent first-order climatic and biophysical controls on wetland CH_4 production, oxidation, and transport. Their relevance to wetland CH_4 emissions is well established in the literature (Knox et al., 2021; Li et al., 2024, 2025; McNicol et al., 2023; Pu et al., 2024; Toet et al., 2011; Yuan et al., 2022, 2024; Zhang et al., 2025). To represent the ecosystem memory effect in response to environmental conditions (Chen et al., 2025), we also included 1-month and 2-month lagged versions of each variable. While introducing short lags would increase collinearity in the machine-learning model, the risk of overfitting is reduced by our model architecture (Section 2.2.2, including feature screening, regularization, subsampling, and early stopping). The data are regridded using bilinear interpolation to 1°x1° resolution to match the wetland CH_4 flux data from the regridded GMB models for further machine-learning modeling.

2.2 XGBoost model

2.2.1 Data assembly and filtering

For each of 350 model-ensemble pairs (35 GMB model estimates x 10 ERA5 ensemble members), we constructed gridded monthly time series pairing GMB data with a set of 35 predictor variables: 11 contemporaneous ERA5 variables, 22 lagged variables (1- and 2-month lags for each of the 11 ERA5 variables), and two seasonal encoding terms:

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$m_{sin}=\sin(2\pi m/12)$, $m_{cos}=\cos(2\pi m/12)$ (Eq.1)

Where:

m=month of year (1-12, 1=January and 12=December);

π=the constant pi;

m_{sin} and m_{cos} are the sine/cosine transforms of m and represent seasonality.

We optimised our analyses to reduce noise and computational burden by retaining only grid cells with substantial wetland emissions, defined as those maintaining a mean flux of $\geq 10^{-15}$ kg $CH_4/m^2/s$ across all GMB model estimates over 2000-2020 (30,000 times lower than the global mean flux).

2.2.2 Model architecture

To systematically reconstruct wetland CH_4 fluxes, we developed a machine-learning pipeline, training independent Extreme Gradient Boosting (XGBoost) models (Chen and Guestrin, 2016) for each grid cell and model-ensemble pair. XGBoost has been applied as a surrogate model for ensemble estimates of GMB BU models with high performance (Zhu et al., 2024). The overall modeling workflow consisted of four stages: (1) data splitting, (2) feature screening, (3) hyperparameter running, and (4) final model fitting and prediction. Each stage is introduced as follows.

Each model-ensemble pair dataset was temporally split into a training set and two testing sets to assess reconstruction capability across different climatic periods at monthly scale. The training data covers January 2003 to December 2017. The testing data covers two periods: January 2000-December 2002 and January 2018-December 2020. Two test windows were used only for out-of-sample evaluation and were not seen during fitting or tuning. The final 24 months of the training period (January 2016 to December 2017) were reserved as a validation set for early stopping and hyperparameter selection. This splitting strategy is to test the model's ability to capture long-term climate variabilities in early/late periods and reduce overfitting.

Prior to XGBoost training, we applied Boruta feature screening (Kursa and Rudnicki, 2010) on the training subset to reduce the effective dimensionality of the predictor set and mitigate overfitting risk arising from the relatively large number of candidate predictors. Boruta augments the predictor matrix with 'shadow' features formed by randomly permuting each predictor across samples, thereby preserving each predictor's marginal distribution while separating its association with CH_4 flux. A random forest regressor is fit to the combined real/shadow feature set and feature importances are compared. Predictor variables were retained if their random forest importance exceeded the maximum shadow-feature importance in more than 20% of 10 iterations (Boruta-style screening), while seasonal terms were always retained. The resulting per-grid cell feature subset was then used for both XGBoost fitting and subsequent reconstructions, ensuring that training and prediction used an identical, cell-specific predictor variable set.

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After the feature screening, we performed a grid search over six hyperparameter configurations (learning rate, maximum tree depth, minimum child weight, subsampling ratio, column subsampling per tree, and L2 regularization) (Table S3). For each configuration, training stopped if the validation Root Mean Squared Error (RMSE) did not improve for 50 consecutive boosting rounds. Once the optimal hyperparameter set was identified for a grid cell, the model was refitted on the full training data using the optimal number of boosting iterations selected during early stopping.

After that, the saved booster was then applied to predict wetland CH_4 flux 2000-2025 (including 2021-2025, which extends beyond the most recent GMB estimates period). All the predicted wetland fluxes (unit of kg $CH_4/m^2/s$) were converted to emissions (Tg CH_4/month) by multiplying the fluxes by grid-cell area (m^2) and total number of seconds (s) in each month.

We recognize that many of these variables are multicollinear, an issue that XGBoost is considered robust against. Nevertheless, we minimized overfitting by constraining model complexity (feature screening, regularized trees with subsampling), using a held-out validation window with early stopping to select the optimal number of boosting iterations, and evaluating performance on temporally separated test periods (2000-2002, 2018-2020) not seen during training. In addition, using the 10 ERA5 ensemble members allows us to propagate reanalysis uncertainty in the predictor fields through the reconstruction.

2.2.3 Model evaluation

Model performance was evaluated for each grid cell per model-ensemble pair by computing the R^2, RMSE, and NRMSE (normalised RMSE, computed by dividing the RMSE by the standard deviation of emissions) values on the withheld test periods. The mean R^2, RMSE and NRMSE of pairs per grid cell are applied to evaluate model performance at the grid cell level (Figure 1). To assess the model reconstruction skills for grid cells with different emission magnitudes, we grouped grid cells into (1) emission percentile bins based on their total wetland CH_4 emissions during 2000-2020 from GMB estimates, (2) 18 geographical regions defined in Saunois et al. (2025) (they are USA, Canada, Central America, Northern South America, Brazil, Southwest South America, Europe, Northern Africa, Equatorial Africa, Southern Africa, Russia, Central Asia, Middle East, China, Korea-Japan, South Asia, Southeast Asia, and Australasia), and (3) five latitude bands: global, Northern Hemisphere (NH) high-latitudes (60°N-90°N), NH mid-latitudes (30°N-60°N), tropics (30°N-30°S), and Southern Hemisphere (SH) extratropics (30°S-90°S). We then compared the mean R^2 and RMSE values of each emission bin (0-10% denotes the highest emitting grid cells) and 18 geographical regions (Figure S1, S2). Monthly mean R^2 and RMSE at five latitude bands over test periods are presented to study the temporal variation of model performance (Figure S3, S4). Detrended monthly global emissions from GMB and predictions were compared to assess the model performance in capturing interannual variability (Figure S5). Detrended anomalies were calculated by fitting a linear least-squares trend to each monthly time series over 2000-2020 and subtracting the corresponding best-fit linear trend from the original series. We generated out-of-sample wetland CH_4 emission predictions (CH_4_pred) for the test periods to compare

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with the emissions from GMB models (CH_4_GMB) (Figure 2). This ensured that the model predicted CH_4_pred based on the trained model without having seen the CH_4_GMB in the test data.

2.3 Emission trend analyses

We quantified spatial patterns, regional aggregates and temporal emission trends in reconstructed wetland CH_4 emissions for 2000-2025, and compared them with corresponding GMB estimates over 2000-2020. We performed analyses for 18 geographic regions and for five latitude bands.

We analyzed recent emission changes by comparing 2021-2025 against 2000-2020 baseline at both grid-cell and regional scales (Figures 3, 4). For Figure 3, gridded monthly emissions were aggregated to per-grid mean emissions for each period (2021-2025 and 2000-2020). Emissions presented in Figure 4 were summed across grid cells within each region. Emission changes between the two periods were calculated as paired differences for each grid cell and model–ensemble pair. Therefore, the reported mean and 95% confidence interval for emission changes reflect the uncertainty of the paired differences themselves. We examined interannual variability and long-term annual emission changes by aggregating monthly emissions to annual totals for both the latitude bands and 18 geographic regions (Figures 5, S6, S8). Ordinary least squares regression was applied to monthly emissions over 2000-2025 to estimate emission trends (Figure 6).

2.4 Uncertainty analysis

Emulator skill uncertainty was evaluated using held-out test period R^2, RMSE, NRMSE and prediction residuals relative to the GMB fields (Sections 2.2.3 and 3.1). Uncertainty in reconstructed wetland CH_4 emissions was estimated from the ensemble spread across all reconstructions generated from the 35 GMB estimates and 10 ERA5 ensemble members. This propagated two sources of uncertainty through the framework: (i) spread in the underlying GMB estimates and (ii) uncertainty in the predictors represented by the ERA5 ensemble. We use 95% confidence interval (CI) to represent this emission uncertainty.

3 Results and Discussion

3.1 Model performance evaluation

We begin by evaluating model performance at the spatial scales most relevant to CH_4 budgeting. Aggregated annual wetland CH_4 emissions at the global and latitudinal-band scales show that the emulator reproduces the overall magnitude, temporal variability, and broad zonal structure of the GMB estimates over 2000-2020 (Figure S6), providing a large-scale context for the finer grid-cell evaluation presented below.

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Across the two test windows (2000-2002 and 2018-2020), the XGBoost model reproduces monthly wetland CH_4 emissions well. Globally, the coefficient of determination R^2 of 0.65 ± 0.003 (mean ± 95% CI, hereafter; Figure 1a), which refers to the mean grid-cell R^2 across all global land cells over the two test periods, with the 95% confidence interval calculated across grid cells. This reported reconstructive skill indicates strong predictability from predictor variables. However, the model performance varies geographically. Mean R^2 is higher in the Southern Hemisphere (SH) than in the Northern Hemisphere (NH) (0.77 ± 0.006 compared with 0.62 ± 0.004, respectively), and higher in the tropics than in the extratropics (0.78 ± 0.004 and 0.58 ± 0.004, respectively). In the top emitting regions, mean R^2 values are generally higher than in regions with lower emissions (Figure S1a). Global RMSE over test periods is 5.49 ± 0.12 $\times 10^{-3}$ Tg CH_4/year (Figure 1b). RMSE for the tropics is about twice that of the RMSE for the extratropics (9.37 ± 0.27 $\times 10^{-3}$ and 3.35 ± 0.08 $\times 10^{-3}$ Tg CH_4/year, respectively). Top emitting regions exhibit higher RMSE, e.g., RMSE doubled for top 10% emitting regions compared with top 10-20% emitting regions (Figure S1b). Regionally, mean R^2 values are highest in Southeast Asia and Korea-Japan (>0.8), and lowest in Russia and Canada (~0.5) (Figure S2a). Mean RMSE is highest in the Americas (Northern South America and Central America), and lowest in the Middle East and Central Asia. Generally, NRMSE is inversely correlated to R^2, indicating a better performance of our model in top emitting regions (Figure 1c). Temporal variation of model performance reveals the lowest R^2 in Jan/Feb in NH high latitudes (Figure S3), and the highest RMSE in July-Sep in the tropics (Figure S4). We further evaluate how these spatial and temporal differences in model skill influence long-term variability in annual emissions in Section 3.3.

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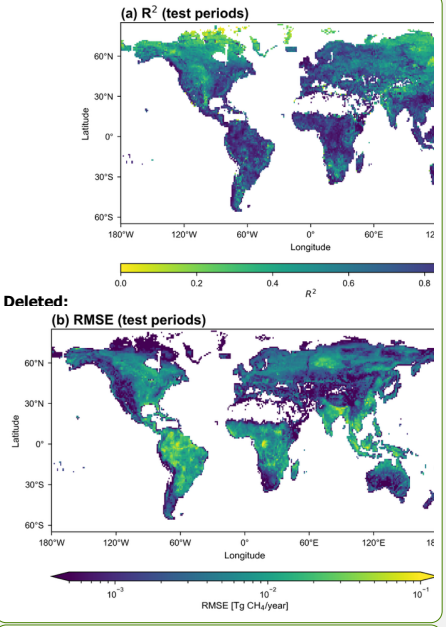

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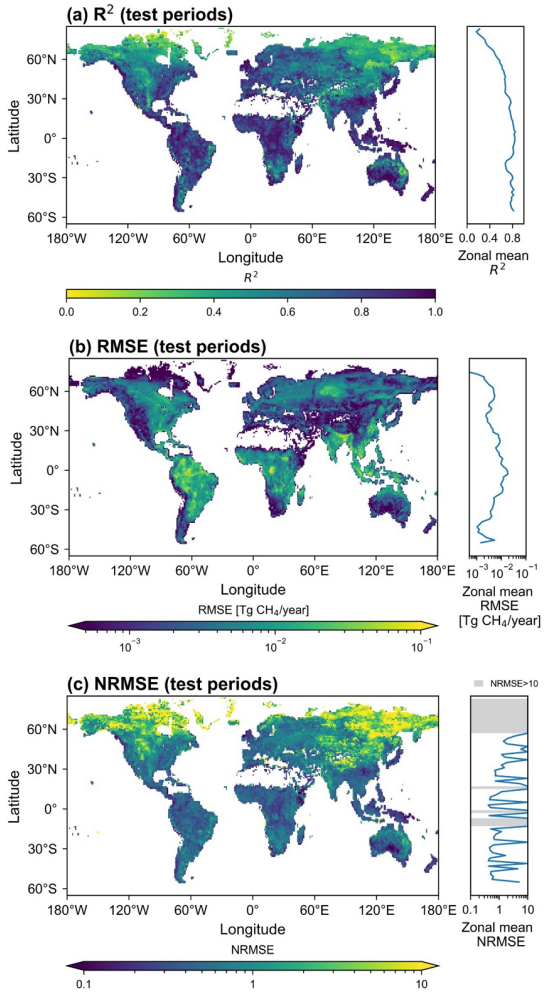

Figure 1. Predictive skill of the XGBoost wetland CH_4 model (per-grid-cell R^2, RMSE, and NRMSE over the test periods). Insets indicate the zonal mean R^2, RMSE and NRMSE in 2° latitude bands.

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To examine predictive uncertainty, we compare wetland CH_4 emissions from the GMB models (CH_4_GMB; Figure 2a) with XGBoost predictions (CH_4_pred; Figure 2b) over the two test periods. Globally, CH_4_pred reproduces both the magnitude and spatial pattern of CH_4_GMB, including the major hotspots concentrated in the tropics, such as Amazon Basin (central-western Amazon floodplains, Pantanal and Moxos plains), Congo Basin, the Sudd wetlands in South Sudan, and Southeast Asia (Sumatra, Borneo, Peninsular Malaysia, and New Guinea), with additional hotspots in NH high latitudes, such as West Siberian Plain and Hudson/James Bay Lowlands.

The differences in emissions (Figure 2c) and the corresponding percent emission differences (Figure 2d) between CH_4_GMB and CH_4_pred highlight spatial biases. During the test periods, the global mean predicted wetland CH_4 emission is 158.5 ± 2.1 Tg CH_4/year, which is 2.3 Tg CH_4/year lower than the global mean of CH_4_GMB (160.7 ± 2.2 Tg CH_4/year), corresponding to a 1.4% underestimation (Figure 2c). This difference is within the estimated uncertainty range of the two global means. Underestimation is strongest over high-emission systems, such as the central-western Amazon Basin, Southeast Asia, Congo Basin and West Siberian Plain (Figure 2c), where our model predicted emissions are lower than GMB estimates, but percent differences are smaller because the denominator (CH_4_GMB) is large. For example, in Southeast Asia, the mean bias is -0.44 Tg CH_4/year, yet the relative bias is only -1.7% given a high regional mean emission (CH_4_GMB of 25.2 Tg CH_4/year). Monthly anomalies in global emissions reveal small disagreements between the predictions and GMB estimates (~±4 Tg CH_4/year) (Figure S5). Largest discrepancies occur in July/August 2020 (underestimation of 15 Tg CH_4/year).

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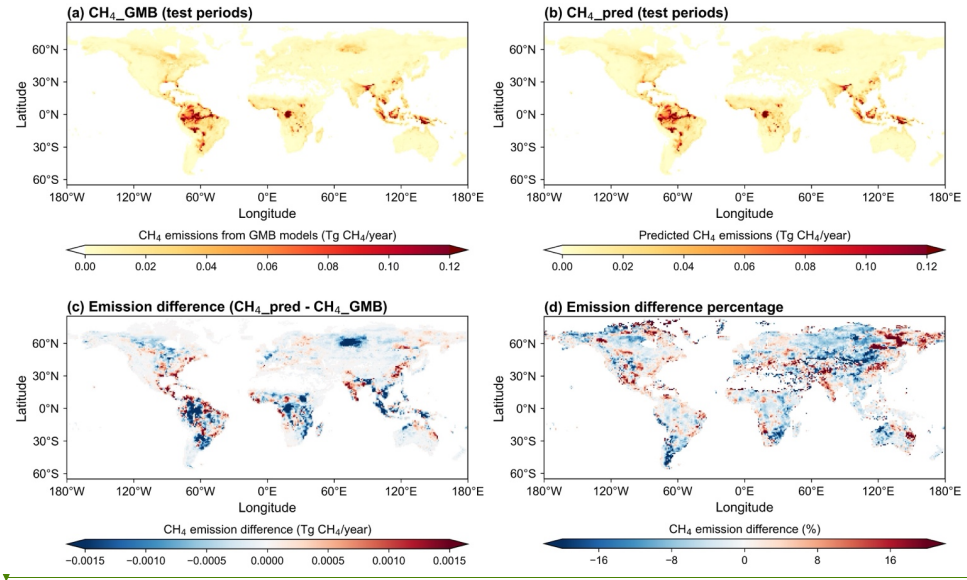

Figure 2. Wetland CH_4 emissions (per 1°x1° grid cell) from GMB models and XGBoost predictions for two test periods combined. (a) Mean CH_4_GMB (GMB emissions) over all test months. (b) Mean CH_4_pred (XGBoost prediction) over all test months. (c) Differences in mean CH_4_GMB and mean CH_4_pred. (d) Percent emission difference between mean CH_4_GMB and mean CH_4_pred. Reds in (c) and (d) indicate higher emissions in CH_4_pred, blues lower emissions.

3.2 Emission trends and predicted anomalies in recent years (2021-2025)

Figure 3 shows spatial maps of global predicted mean wetland CH_4 emissions for 2021-2025 (Figure 3a) and for 2000-2020 (Figure 3b). The 2021-2025 mean emission is 157.8 ± 2.4 Tg CH_4/year (z= 3.72, with the z-score indicating how far each region's mean emission estimate deviates from the cross-region mean, measured in standard deviation units), representing a non-significant increase of 0.05 Tg CH_4/year relative to the 2000-2020 mean. However, changes are more pronounced across latitudes. NH mid- and high-latitude wetland emissions increase during 2021-2025 by 0.76 ± 0.07 (from 29.9 ± 1.1 in 2000-2020 to 30.7 ± 1.2 in 2021-2025, z=2.21) and 0.35 ± 0.03 (from 11.0 ± 0.7 to 11.4 ± 0.7 Tg CH_4/year, z=1.01), respectively, compared with 2000-2020 emissions. These changes correspond to ~2.5% and ~0.8% increase. The continued growth in these regions in 2021-2025 (relative to 2000-

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2020), together with increases reported for 2010-2018 relative to 1981-1989 (Feron et al., 2024), suggests an increasing role of boreal and temperate wetlands in recent decades of global warming. In contrast, emissions decrease in the tropics and SH extratropics during 2021-2025. The tropics exhibit a 0.95 ± 0.19 Tg CH_4/year emission decrease (z=-2.81), from 113.9 ± 2.5 (2000-2020) to 112.9 ± 2.5 Tg CH_4/year (2021-2025), while SH extratropics emissions decline by 3.5% (-0.11 ± 0.02 Tg CH_4/year, z=-0.34)) from 3.1 ± 0.1 (2000-2020) to 3.0 ± 0.1 Tg CH_4/year (2021-2025).

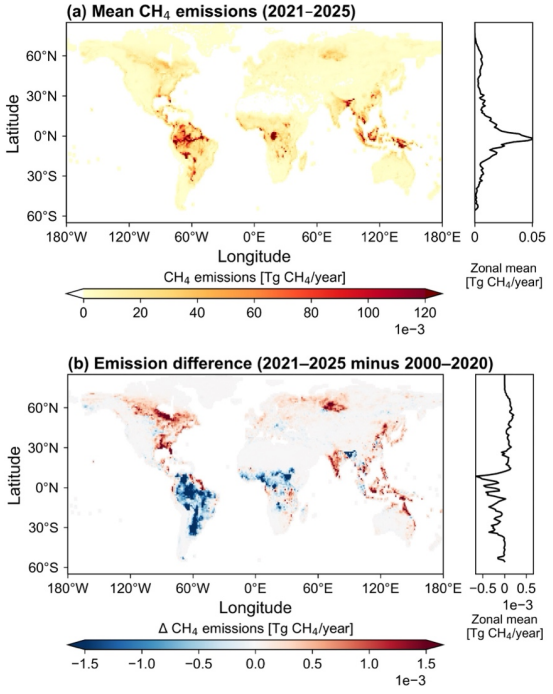

Figure 3. (a) Mean XGBoost model predicted wetland CH_4 emissions for 2021-2025. (b) Difference between mean emissions in 2021-2025 and 2000-2020. Reds indicate higher emissions in 2021-2025, blues lower. Insets indicate the zonal mean in 2° latitude bands.

We further assess regional wetland CH_4 emission changes in 2021-2025 compared with 2000-2020 across 18 regions using XGBoost predictions (Figure 4) across 18 regions. Four regions show no significant change (p≥0.05 from two-sided paired significance test, labelled "ns"). Among the 14

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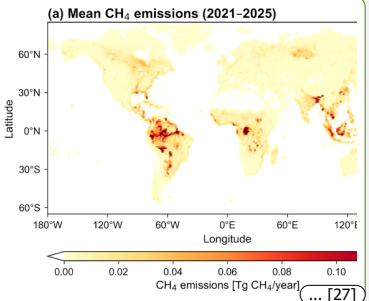

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regions with significant changes, nine regions show increases and five show decreases (Northern South America, Brazil, Southwest South America, Northern Africa, and Equatorial Africa). Notably, all regions with decreases are located in South America and Africa. The largest emission changes (2021-2025 relative to 2000-2020) occur in Brazil (-0.59 ± 0.09 Tg CH_4/year), Canada (+0.46 ± 0.05 Tg CH_4/year), Southwest South America (-0.43 ± 0.07 Tg CH_4/year), Russia (+0.31 ± 0.03 Tg CH_4/year), Southeast Asia (+0.22 ± 0.03 Tg CH_4/year) and Equatorial Africa (-0.21 ± 0.04 Tg CH_4/year). The pronounced decline over Brazil possibly reflects the exceptional drought conditions in the Amazon basin during 2022-2024, including record-low river levels and anomalously warm, dry conditions (Espinoza et al., 2024). Such hydroclimatic drying reduces floodplain inundation and lowers water tables, which suppresses anaerobic conditions and thereby limits wetland methane production and emissions (Cui et al., 2024).

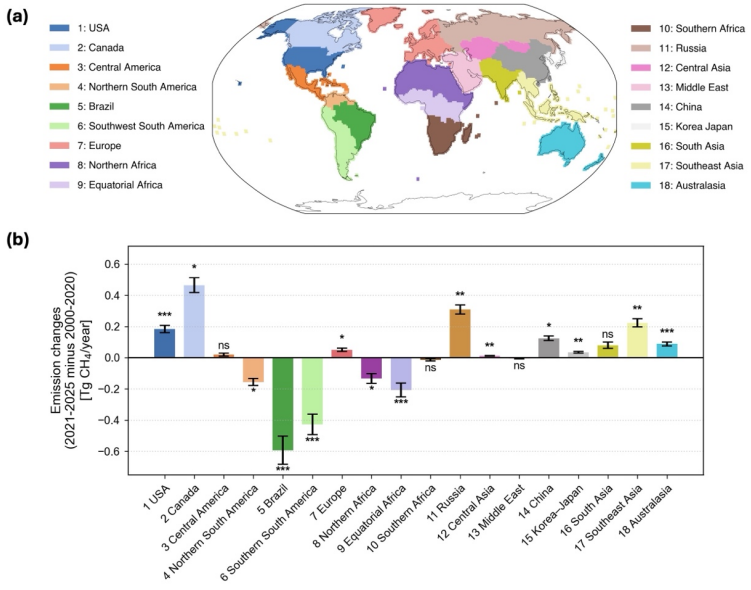
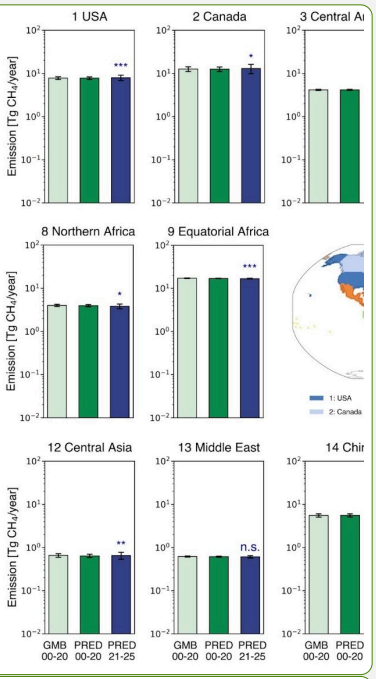


Figure 4. (a) Spatial distribution of the 18 regions; (b) Regional wetland CH_4 emission changes during 2021-2025 relative to 2000-2020 across 18 regions using XGBoost modeled emissions. Bars show regional mean emission changes (annual mean in 2021-2025 minus that in 2000-2020). Error bars indicate 95% CI. A two-sided paired test was applied for each region to assess whether the 2021-2025 mean emission differs from the 2000-2020 mean. Regions with p≥0.05 are labelled "ns". Asterisks above bars denote significance: *** for p < 0.001, ** for p < 0.01, * for p < 0.05.



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3.3 26-year emission trends and anomalies (2000-2025)

The emission predictions enable an assessment of long-term trends and interannual variability in wetland emissions over 2000-2025. Annual emission time series show pronounced peaks in 2011 and 2016/2017 in the tropics, NH mid-latitudes, and globally (Figure S6). Monthly emission anomalies in these regions (Figure S7) indicate that the annual maxima are driven by sustained positive anomalies during Aug 2010-May 2011 and Aug 2016-Dec 2017, which coincide with La Niña conditions. This alignment is consistent with prior studies reporting La Niña-driven wetting and expanded inundation that enhance tropical wetland CH_4 emissions (Hodson et al., 2011; Lin et al., 2024; Murguia-Flores et al., 2023; Zhang et al., 2020, 2018; Zhu et al., 2017), and with the broader evidence that CH_4 growth in 2020-2022 coincided with an unusual persistent La Niña event.

In the 2020s, global and tropical annual emissions reach a minimum in 2023 (Figures S6, S7), coincident with a strong El Niño event. This is supported by two recent studies reporting a sharp decline in wetland emissions in South America (Ciais et al., 2026) and Amazonia (Quinn et al., 2025) in 2023 linked to El Niño-related drought. Our results further suggest that Brazil and Southern South America are the key contributors, and both rank among the top six emitting regions globally (Figure 5). Annual emissions in Brazil decline steadily across the study period, from 24.8 ± 0.9 Tg CH_4/year in 2000 to 23.3 ± 0.9 Tg CH_4/year in 2025, with the lowest emission in 2024 (22.8 ± 0.9 Tg CH_4/year) (Figure 5). Emissions in Southern South America are relatively stable during 2000-2014, decrease markedly during 2014-2023 (from 17.2 ± 0.6 to 15.6 ± 0.6 Tg CH_4/year), and show a modest recovery in 2024/2025 (Figure 5).

The remaining four top emitting regions show contrasting behavior. Since 2019, Southeast Asia, Canada and Russia show increasing emissions, whereas Equatorial Africa emissions decrease slightly (Figure 5). Southeast Asian emissions fluctuate around 24.7 ± 0.5 Tg CH_4/year over 2000-2025, with a minimum in 2015 (23.8 ± 0.7 Tg CH_4/year) and a maximum in 2017 (25.6 ± 0.7 Tg CH_4/year). Emissions in Canada are comparatively stable during 2005-2015, while variability is larger before 2005 and after 2015: the amplitude reaches ~1.5 Tg CH_4/year in these periods, approximately double that during 2005-2015 (~0.8 Tg CH_4/year). Equatorial African emissions remain 17.0 ± 0.2 Tg CH_4/year during 2000-2025, despite two notable declines during 2002-2005 (-0.5 Tg CH_4/year) and 2019-2025 (-0.6 Tg CH_4/year).

For regions showing increasing emissions in our XGBoost predictions (e.g., Southeast Asia, Canada, and Russia), these changes are likely driven by wetter and/or warmer conditions that enhance inundation, substrate supply, and CH_4 production. For instance, observed increases in Boreal-Arctic wetland CH_4 emissions have been linked to warming and enhanced ecosystem productivity (Yuan et al., 2024), while inversion-based studies highlight the role of La Niña-driven inundation in enhancing emissions from Equatorial Asia (Qu et al., 2024; Lin et al., 2024; Feng et al., 2022). However, our emulator predictions diverge from recent atmospheric inversion literature regarding African wetland emissions. Multiple top-down studies indicate that enhanced emissions from tropical Africa contributed substantially to the global CH_4 surge during 2020-2024 (Balasus et al., 2026; Qu et al., 2024). In

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contrast, our climate-driven emulator predicts that emissions across Northern, Equatorial, and Southern Africa remained relatively stable or slightly decreased during similar time periods (Figures 4, 5, S8). We attribute this discrepancy to weaker performance and regional precipitation biases of ERA5 data over parts of Africa in capturing complex African hydrological dynamics (Gebrechorkos et al., 2024).

Regional emission time series (Figures 5, S8) help explain the pronounced disagreement in 2019-2020 for global and tropical emissions, where predictions are substantially lower than the GMB estimates (Figure S6). For 2019, the prediction mean emissions are 5.5 Tg CH_4/year below the GMB estimates for global emissions (~3.5% underestimation), while emission changes in Equatorial Africa, Brazil, USA, Southern Africa, and Southern South America jointly account for 4.3 Tg CH_4/year of this discrepancy. In 2020, the modeled mean global emissions are 9.8 Tg CH_4/year lower than GMB (~6% underestimation). The discrepancies in the tropics explain 9.2 Tg CH_4/year of the underestimation, including major regions such as Eastern Africa, Southern South America, Southeast Asia, Brazil, and Northern Africa. Notably, this disagreement in 2019-2020 is not evident in other latitude bands and does not persist outside these years. We cautiously interpret this mismatch as the limitations of the emulator and uncertainty in the underlying GMB ensemble. The climate predictors used in this study can only indirectly capture ENSO (El Niño-Southern Oscillation) effects. Although the training period includes several ENSO events, the unusual persistence of the 2020-2023 triple-dip La Niña may still be underrepresented. In addition, most TD estimates from GMB use a similar prescribed OH (hydroxyl radical) distribution, possibly causing OH-related uncertainty to be underrepresented, particularly during anomalous periods such as COVID.

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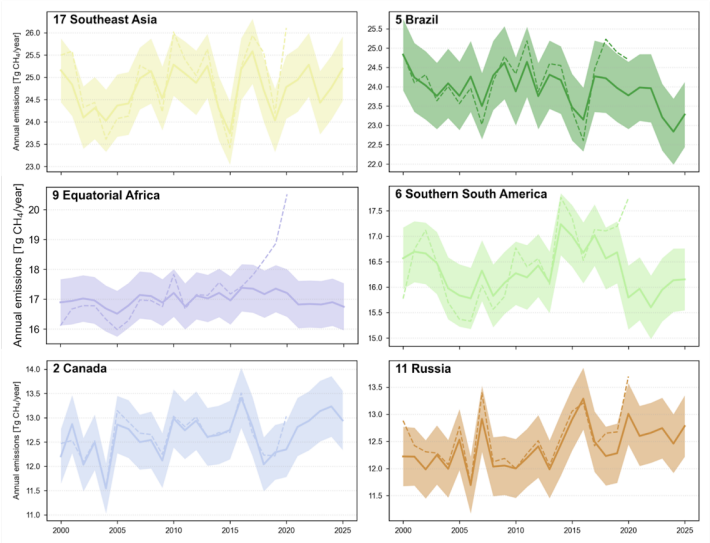
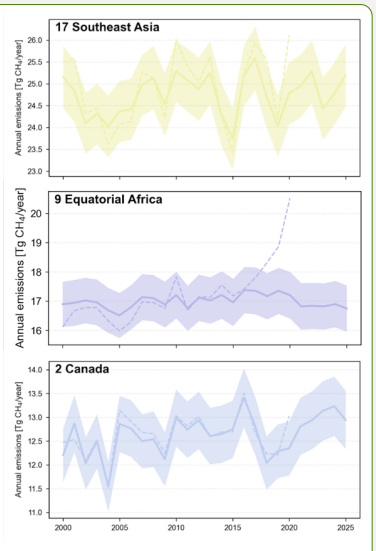


Figure 5. Annual wetland CH_4 emissions for the six top-emitting regions. Solid lines show model predictions of annual mean emissions, and shaded areas indicate 95% CI. Dashed lines show GMB emission estimates (mean). Panel titles indicate the region code and region name, which correspond to the region definitions shown in Figure 4.

Finally, we examine trends in predicted monthly wetland CH_4 emissions across five latitude bands (Figure 6a) and 12 calendar-month categories (Figure 6b). Figure 6a compares these trends across three time periods (2000-2025, 2021-2025, and 2000-2020). Figure 6b shows the trends by 12 calendar-month categories for the 2000-2025 period. We summed predicted monthly emissions across all grid cells within each latitude band to form regional monthly time series, then computed monthly emission anomalies by subtracting the 2000-2025 climatological mean. We then estimated trends (Tg CH_4/year) using ordinary least squares linear regression of the anomaly time series. We calculated a z-score for each region and period by dividing the estimated trend by its standard error to assess the statistical significance of the trends.

The global wetland CH_4 emission trend is three times lower in 2021-2025 compared with the 2000-2020 period (0.05 ± 0.28 and 0.15 ± 0.05 Tg CH_4/year, respectively) (Figure 6a). However, the magnitude of changes in emission trends in NH mid-latitudes, the tropics and SH extratropics is greater in 2021-2025 than 2000-2020 level. Emission trends in NH mid-latitudes increased three times from 0.05 ± 0.02 Tg CH_4/year (z=4.94) in 2000-2020 to 0.16 ± 0.12 Tg CH_4/year (z=2.49) during 2021-2025. The tropics exhibit a reversal of their emission trends from positive in 2000-2020 (0.07 ± 0.03 Tg CH_4/year,



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z=3.95) to negative in 2021-2025 (-0.18 ± 0.24 Tg CH_4/year, z=-1.45). Emission trends in SH extratropics increased from nearly zero in 2000-2020 (-8.74 x10^{-4} ± 5.18 x10^{-3} Tg CH_4/year, z=-0.33) to higher than the global average in 2021-2025 (0.04 ± 0.02 Tg CH_4/year, z=3.30).

Evaluating emission trends over individual months (Figure 6b), we find that the global trend peaks in late boreal summer, with the largest increases in August and September (0.15 ± 0.18 and 0.21 ± 0.13 Tg CH_4/year, respectively). The strongest seasonal intensification occurs in NH mid-latitudes during June-September (0.09-0.12 Tg CH_4/year), consistent with the dominant contribution from boreal growing-season emissions. For both NH mid- and high-latitudes, growth rates are systematically higher during May-October (growing season) than during November-April. In the tropics, monthly growth reaches its highest in Oct-Dec and lowest in Jun-Aug (negative growth). SH extratropics show near-zero growth throughout the year (slightly negative).

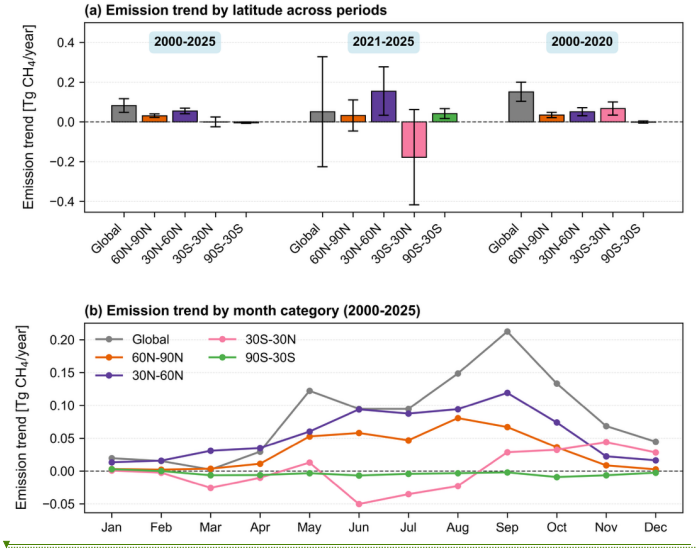

Figure 6. Regional and seasonal trends (Tg CH_4/year) in predicted wetland CH_4 emissions. (a) Trends across five latitude bands across three time periods (2000-2025, 2021-2025, and 2000-2020). (b) Trends for each calendar month over 2000-2025. Positive trends indicate an increase in CH_4 emissions, while negative values indicate a decrease. Error bars indicate 95% CI.

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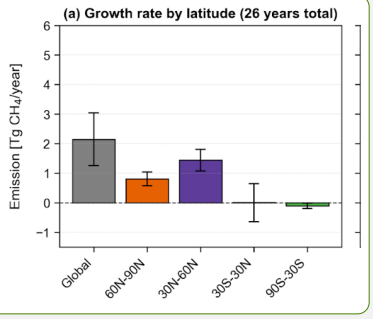

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3.4 Limitation

A limitation concerns source attribution in the TD estimates used for training. Although the TD inversions adopted BU wetland priors that exclude inland freshwater systems, the coarse spatial resolution of TD inversions can still limit the separation of natural vegetated wetland emissions from nearby open-water or inland-freshwater emissions in mixed floodplain environments. This uncertainty cannot be quantified explicitly in the present framework because no independent source-resolved posterior diagnostic is available. Future progress is likely to come from higher-resolution atmospheric inverse frameworks. Two recent studies performed 25 km resolution atmospheric inversions using satellite observations in regional and global applications and improved spatial attribution (East et al., 2025; Hancock et al., 2025).

Another limitation is that our framework predicts wetland emissions from climate predictors only. ENSO-related impacts can be captured only indirectly, while dynamic inundation and atmospheric chemistry perturbations are outside the predictor space of the model.

4 Conclusions

This study presents a framework to extend global natural wetland CH_4 emissions from 2000 through 2025 by applying a machine-learning emulator to provide a lower-latency release of GMB estimates to estimate wetland emissions and is intended for routine updates. The framework reproduces both spatial patterns and interannual variability of the GMB estimates. The extended wetland CH_4 emission record shows that the magnitude of emission changes in 2021-2025 is substantially larger than the long-term trend over 2000-2020, even though the global mean emission change is relatively small.

A key emerging feature is accelerated emissions estimated in the NH mid-latitudes and high-latitudes. Seasonal trend analyses indicate that these increases are concentrated in the warm season (May - Oct), consistent with an amplification of the seasonal cycle and a strengthening of late-summer growth, highlighting Northern mid- and high-latitude wetlands as regions where intensifying trends are most evident in recent years.

In contrast, the tropics show a reversal in emission trend in the post-2020s, from an emission increase during 2000-2020 to a decline in 2021-2025. Nevertheless, the tropics still dominate interannual variability of global wetland emissions and can strongly influence the global CH_4 budget through episodic hydroclimate extremes.

More broadly, this emulator framework may also apply to other low-latency Earth system applications beyond wetland CH_4 emissions, particularly for environments and pollutants that can be constrained by routinely updated predictor fields. At the same time, the present results point to several opportunities for

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improvement, including better representation of dynamic inundation and improved separation of wetland and inland-freshwater emissions.

Overall, this work proposes an operational, low-latency emulator framework that provides a scalable pathway to track wetland CH_4 emissions in response to climate anomalies by leveraging routinely updated Earth observation reanalysis data.

Data Availability

ERA5 monthly averaged data are available at Copernicus Climate Change Service Climate Data Store at https://cds.climate.copernicus.eu/datasets/reanalysis-era5-single-levels-monthly-means?tab=overview (DOI:10.24381/cds.f17050d7) (Accessed on 22-Jan-2026) (Copernicus Climate Change Service, 2023). The global natural wetland methane emission dataset generated from this study is publicly available at https://doi.org/10.5281/zenodo.18870108 (Li et al., 2026a).

Competing Interests

At least one of the (co-)authors is a member of the editorial board of Earth System Science Data.

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