

GLBD-FED: A Global first-hand in-situ daily temperature dataset preferentially with a 0000-2400 UTC 24-hour window (1981-2024)

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1. Abstract

Large amounts of sub-daily temperature data are shared globally through the Global Telecommunication System (GTS) in near-real time and through international data exchanges. However, converting these data into a global daily temperature dataset with a uniform definition—especially for daily maximum (Tmax) and minimum (Tmin) temperatures—has proven challenging due to the independent observation schedules across the world. To address this issue, we developed a new method that decomposes sub-daily Tmax and Tmin records from the Integrated Surface Database (ISD) into finer intervals, subsequently reaggregating them into daily Tmax and Tmin based on a prospective 0000–2400 UTC physical window. This new method increased the global daily Tmax and Tmin data counts by 64% and 45%, respectively, compared to the conventional calculation method, which relied on either two consecutive Tmax/Tmin records over 12 hours or a single Tmax/Tmin record over 24 hours. The Global Land Base Dataset-First Estimate Daily Data (GLBD-FED) was established for the period from 1981 to 2024, following corrections for misrecorded Tmax and Tmin and quality control. GLBD-FED includes daily maximum (Tmax), average (Tave), and minimum temperature (Tmin) from approximately 17,000 global sites, with daily data amounts reaching about 10,000 entries per day in the current decade. When compared to the Global Summary of the Day (GSOD) dataset, GLBD-FED exhibits less extreme daily values over the last 40 years, showing a slightly lower daily Tmax (around -0.3°C) and a higher daily Tmin (around +0.3°C), with a nearly

identical daily Tave (approximately +0.1°C). These systematic differences arise from multiple sources, including: (1) UTC-boundary double-counting artifacts in GSOD (which alone introduced an average bias of +0.28°C in Tmax and -0.85°C in Tmin across ~18 million and ~11 million duplicated records, respectively); (2) differing sub-daily data source preferences (SYNOPTIC vs. METAR reports); and (3) divergent date boundary treatment methods. The database and associated data can be found at <https://doi.org/10.5281/zenodo.17895292> (Yang et al., 2025).

2. Introduction

Near-surface air temperature is one of the most fundamental and core observational elements for reflecting global climate change (IPCC, 2021). In the context of global warming, obtaining long-term, high-spatial-coverage, and quality-controlled global daily instrumental surface temperature data (primarily including daily maximum, minimum, and mean temperatures) plays an irreplaceable role in climate change detection, attribution analysis, and the evaluation of weather and climate models (Alexander et al., 2006). Such high-temporal-resolution observational data not only reflects the mean state characteristics of the climate system but also serves as the cornerstone for characterizing high-frequency weather fluctuations and climate variability.

In operational practices and scientific research, daily temperature data (especially daily maximum and minimum temperatures) are an absolute prerequisite for monitoring and assessing extreme weather and climate events. For instance, a series of core extreme climate indices defined by the WMO Expert Team on Climate Change Detection and Indices (ETCCDI)—such as the Warm Spell Duration Indicator (WSDI), Frost Days (FD0), and Extreme Maximum Temperature (TXx)—must be calculated based on continuous, high-quality daily temperature data (Alexander et al., 2006; Donat et al., 2013; Dunn et al., 2020; Zhang et al., 2011). These derivative indices not only provide indispensable observational evidence for successive assessment reports by the Intergovernmental Panel on Climate Change (IPCC) but are also widely applied in the real-time monitoring of severe weather and early warning systems (IPCC, 2021).

While numerous global observational datasets exist to support climate research, there is a pronounced scarcity of daily global temperature datasets based purely on in situ station data. Historically, the development of global observational products has been exceptionally robust for precipitation, flourishing through both dense in situ gauge-based gridded analyses (e.g., GPCC; Becker et al., 2013) and multi-source merged products incorporating satellite estimates (e.g., MSWEP; Beck et al., 2019). In contrast, the available landscape for global daily temperature is much narrower. Existing prominent temperature datasets serve vital but distinct purposes: GHCN-Daily (Menne et al., 2012) acts as the world's most comprehensive integrated database of daily climate summaries, while datasets like Berkeley Earth (BEST; Rohde et al., 2013) and HadGHCND (Caesar et al., 2006) provide spatially interpolated gridded products for climate analysis. This scarcity leaves a critical gap for a purely station-based, high-fidelity daily temperature dataset that can leverage the improved global accessibility of sub-daily observations to enforce strict temporal uniformity.

High-quality daily temperature datasets generally fall into two distinct methodological approaches: comprehensive integrated databases (such as GHCND) designed to assimilate reports from diverse national networks, and first-hand Synoptic datasets designed for rapid global monitoring. While integrated databases are essential for climatology and provide near-real-time updates in many regions, they inherently preserve varying local daily measurement periods (e.g., local morning versus midnight-to-midnight), which introduces Time of Observation Biases (TOB). Conversely, legacy near-real-time datasets (such as GSOD) attempt to standardize time but exhibit severe temporal aggregation artifacts—such as misallocating sub-daily extremes to incorrect calendar days due to strict UTC boundary constraints (the specific mechanisms and quantitative impacts of these challenges are discussed in detail in Section 5.3). These algorithmic characteristics introduce artificial noise that affects rapid extreme weather analysis. Furthermore, a critical gap remains: strict temporal and methodological uniformity. National and regional benchmark datasets frequently employ diverse definitions of a 'daily' period and utilize disparate temporal aggregation algorithms. This lack of standardization introduces well-documented Time of Observation Biases (TOB; Karl et al., 1986). When researchers attempt to combine these localized high-quality datasets for global monitoring, these

methodological differences inevitably create artificial discontinuities ('seams') across national borders.

The core scientific significance of GLBD-FED lies in its ability to eliminate these methodological borders. By applying a single, unified algorithmic framework directly to first-hand sub-daily SYNOPTIC reports across all global regions simultaneously, GLBD-FED enforces a universal physical 24-hour window (0000 to 2400 UTC). This strict temporal alignment is irreplaceable for advanced meteorological applications, particularly in the verification of global Numerical Weather Prediction (NWP) model outputs. Modern NWP models output daily forecast summaries based on standardized UTC cycles; validating these outputs against heterogeneous regional datasets introduces severe temporal mismatch errors (Haiden et al., 2018). By aligning with the World Meteorological Organization's standard for SYNOPTIC uniformity (WMO, 2017), GLBD-FED provides a seamless, time-aligned 'first estimate' ground truth. Therefore, the primary focus of GLBD-FED is to serve as a high-fidelity, global, near-real-time, first-hand daily dataset that supports reliable, large-scale extreme event monitoring and serves as temporally consistent raw material for future benchmark homogenization.

Accordingly, this study aims to produce GLBD-FED, a global first-hand daily temperature dataset (1981–2024) that preferentially aligns maximum, minimum, and average temperatures to a unified 0000–2400 UTC 24-hour window. To achieve this, we developed new temporal reconstruction algorithms designed to optimally utilize irregularly timed sub-daily measurements from the ISD. By harmonizing these records under a strict temporal framework, our methodology effectively mitigates Time of Observation Biases (TOB) and artificial geopolitical border effects. The data sources and methodology—including automated error correction and sliding-window fallback mechanisms—are detailed in Sections 3 and 4. This is followed in Section 5 by a comparative evaluation against a legacy dataset (GSOD), which quantitatively highlights the resolution of systematic temporal attribution biases. Finally, conclusions are presented in Section 6.

3. Data Sources

3.1 Near-real-time sub-daily meteorological measurements

Hourly surface temperature observations spanning 1981–2024 were compiled through the Integrated Surface Database (ISD), a global database that consists of hourly and SYNOPTIC surface observations compiled from numerous sources into a single common ASCII format and common data model (Smith et al., 2011). It is important to note that NOAA formally terminated updates to the ISD in 2025, transitioning to the new Global Historical Climatology Network hourly (GHCNh) framework (Menne et al., 2025). However, our evaluation indicates that GHCNh currently lacks several critical meteorological elements present in the legacy ISD—specifically the 12-hour and 24-hour temperature extremes. Given that these explicitly reported extremes are essential for our algorithm, the historical ISD remains the only viable source for reconstructing the high-fidelity 1981–2024 baseline presented in this dataset.

The ISD-derived temperature parameters encompass five specific temporal components: discrete sub-daily instantaneous temperatures (which are utilized to derive the daily mean temperature, denoted as Tave), 24-hour explicitly reported maxima and minima (Tmax-24h and Tmin-24h), and 12-hour explicitly reported maxima and minima (Tmax-12h and Tmin-12h).

Fig 1 visualizes the global distribution of sub-daily temperature data volumes across the 24-hour UTC cycle spanning the period from 1981 to 2024. The analysis reveals striking temporal discrepancies among different temperature parameters. For the discrete hourly temperature observations (utilized to derive Tave), the data volume is continuously distributed across all hours, characterized by a highly robust multi-peak pattern. The primary peaks align perfectly with the standard 6-hourly SYNOPTIC times (0000, 0600, 1200, and 1800 UTC), complemented by secondary peaks at the intermediate 3-hourly intervals (0300, 0900, 1500, and 2100 UTC). This dense and temporally consistent distribution provides a solid foundation for calculating highly representative daily mean temperatures. Analysis of the discrete sub-daily reporting frequency reveals that the global network primarily operates on a combination of 3-hourly and 6-hourly cadences, manifesting as two dominant 6-hourly regimes offset by 3 hours: 0000/0600/1200/1800 UTC and 0300/0900/1500/2100 UTC.

In stark contrast, the explicitly reported extreme temperatures (Tmax and Tmin) exhibit extreme temporal concentration. The 24-hour extremes (Tmax-24h and

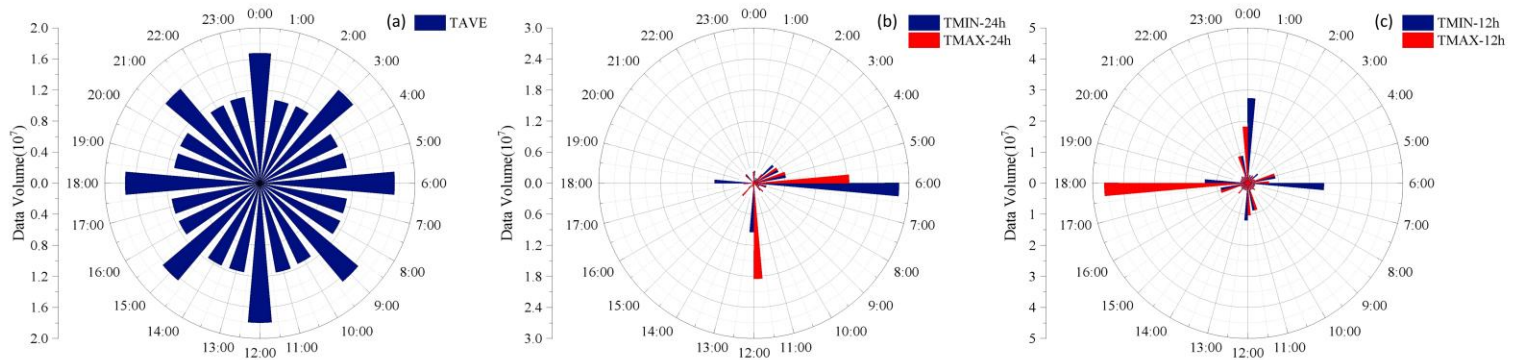
Tmin-24h) are overwhelmingly anchored at just two specific reporting times: 0600 and 1800 UTC. Similarly, the 12-hour extremes present a highly asymmetric reporting pattern, reaching absolute volumetric peaks at 0600 UTC (for Tmin-12h) and 1800 UTC (for Tmax-12h). It is important to clarify that aligning 1800 UTC with daytime maxima and 0600 UTC with nighttime minima is physically applicable only for regions near the Prime Meridian (0° longitude). For the vast majority of the globe, local solar heating and cooling cycles do not coincide with these specific UTC hours. Therefore, this overwhelming global concentration is not a reflection of natural diurnal cycles, but rather an administrative artifact of international SYNOPTIC data exchange protocols. It should be noted that the overwhelming concentration of explicit 24-hour extreme reports at specific hours (e.g., 0600 and 1800 UTC) is primarily an artifact of international data exchange standards, rather than a reflection of the longitudinal/geographic distribution of stations. WMO SYNOPTIC reporting protocols require member states to aggregate and submit extreme records at standardized global synoptic hours (WMO, 2015). Consequently, the ISD archive structurally 'anchors' the vast majority of explicit global extreme reports to these specific UTC timestamps.

These distinct structural characteristics explicitly demonstrate that while hourly observations offer continuous sub-daily coverage, explicit extreme reports are highly sparse outside of a few specific synoptic hours. This temporal fragmentation structurally mandates and quantitatively justifies the necessity of our secondary fallback strategy, which utilizes the high-frequency hourly observations to robustly reconstruct daily extremes when explicit records are absent.

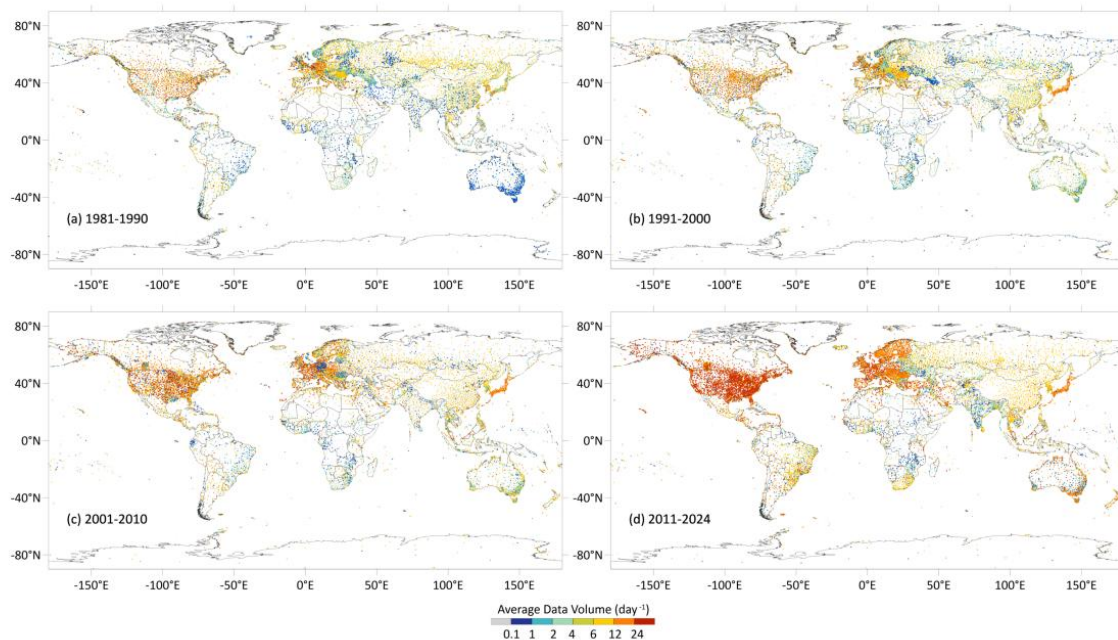
The multi-panel maps in Fig 2 to Fig 6 illustrate the spatiotemporal evolution of sub-daily temperature data volumes from 1981 to 2024. The analysis reveals a striking contrast in data availability: while discrete hourly temperatures exhibit continuous and stable growth in spatial coverage and reporting frequency over the four decades, explicitly reported extremes (12-hour and 24-hour Tmax/Tmin) suffer from severe geographic fragmentation and decadal volatility, notably experiencing a pronounced global decline between 1990 and 2010. These stark spatiotemporal discrepancies visually highlight the limitations of relying exclusively on explicitly reported extremes and quantitatively justify the necessity of utilizing high-density

189 hourly data as a secondary fallback strategy.

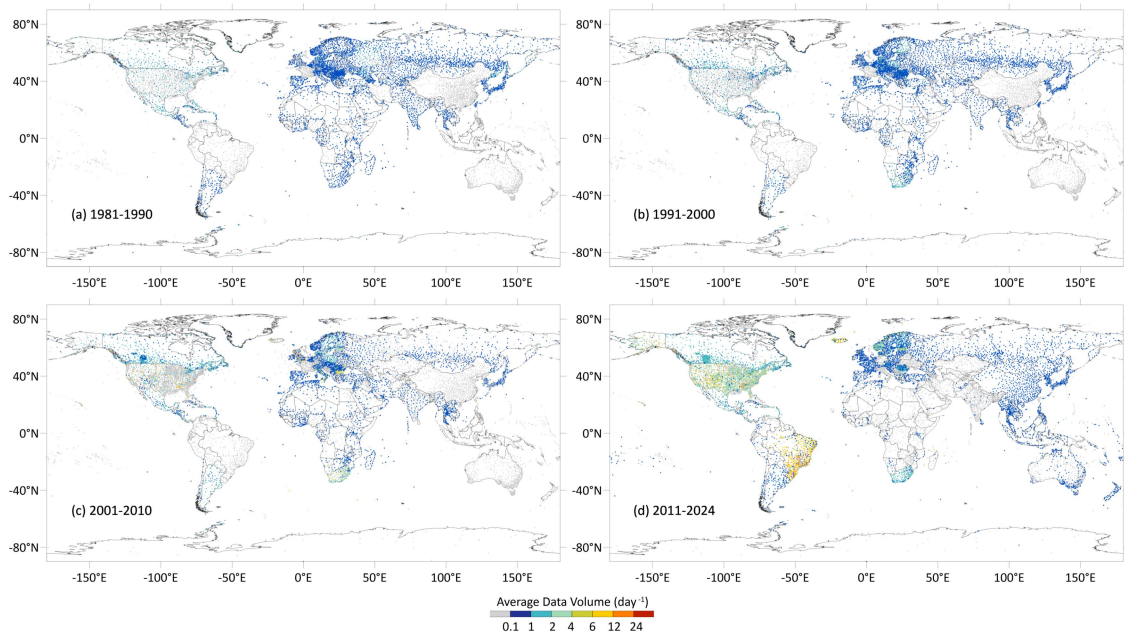
190 In summary, sub-daily observations exhibit irregular observation intervals compared
 191 to Tave. Moreover, Tmax and Tmin do not always appear in pairs as expected,
 192 indicating that date shifts and inhomogeneities may not only exist between regions but
 193 also between the raw Tmax and Tmin records themselves.



194 Figure 1 The distribution of sub-daily temperature data amounts at each o'clock during 1981–2024.
 195 (Note: In panels b and c, the red and blue bars are purposefully offset slightly for visual clarity to
 196 prevent overlap, though they represent records reported at the exact same synoptic hours.)

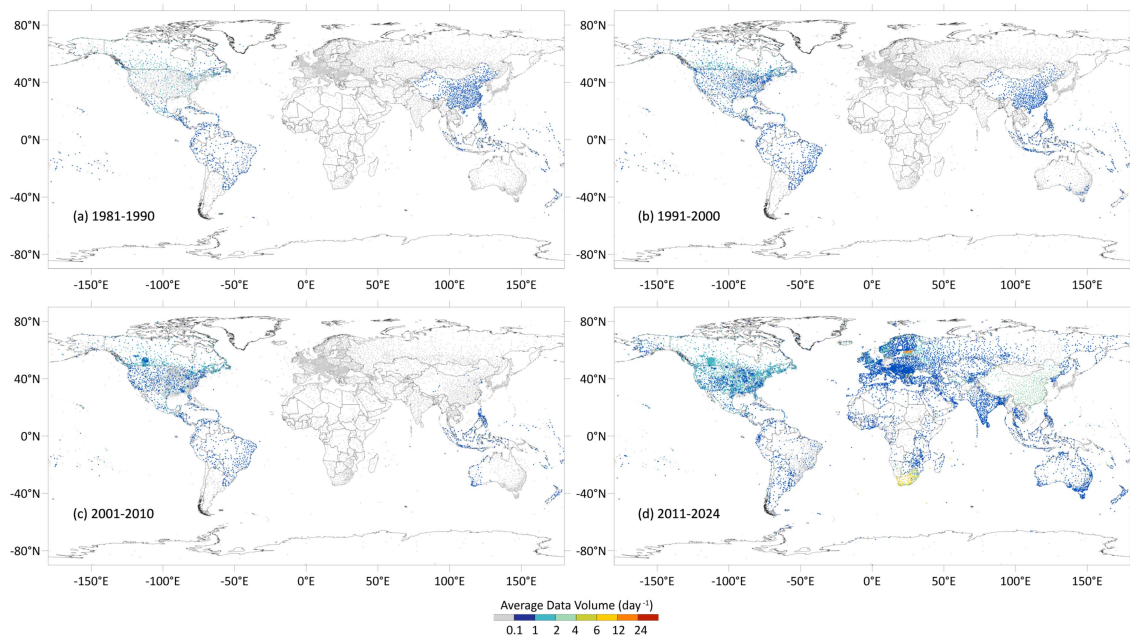


197 Figure 2 Spatial distributions of Hourly average data volume per day for hourly Tave from ISD. Panel
 198 a,b,c,d stand for the results 1981-1990, 1991-2000, 2001-2010, 2011-2024. Spatial distributions of
 199 hourly average data volume per day for hourly Tave from ISD. Panels a, b, c, d represent 1981 – 1990,
 200 1991 – 2000, 2001 – 2010, 2011 – 2024.



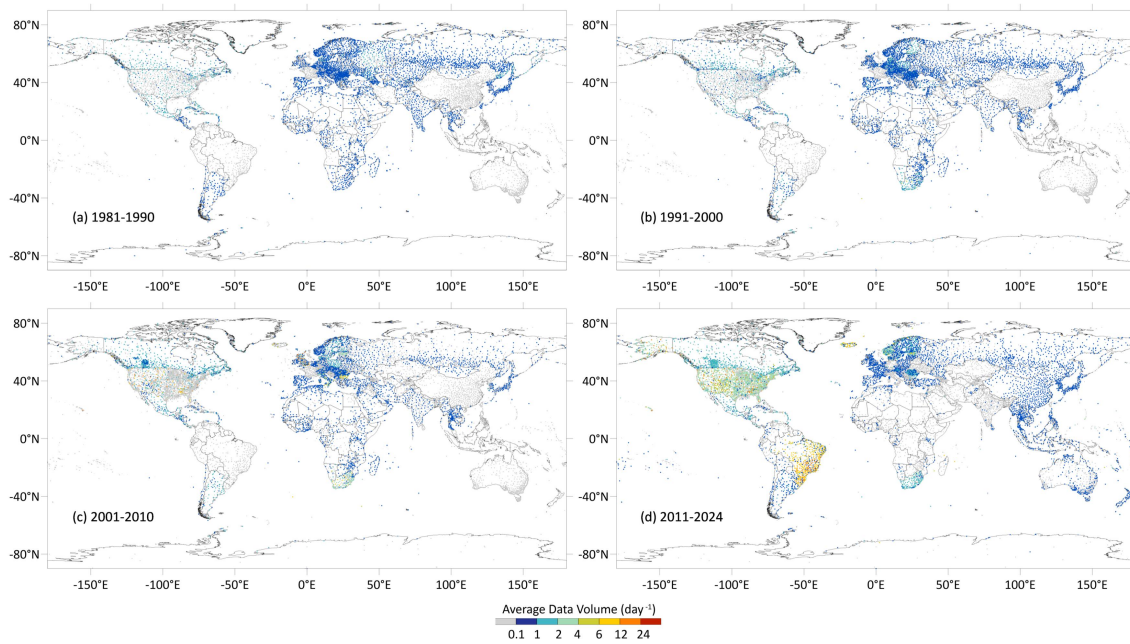
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Figure 3 Similar to Figure 1, but for Tmax-12h



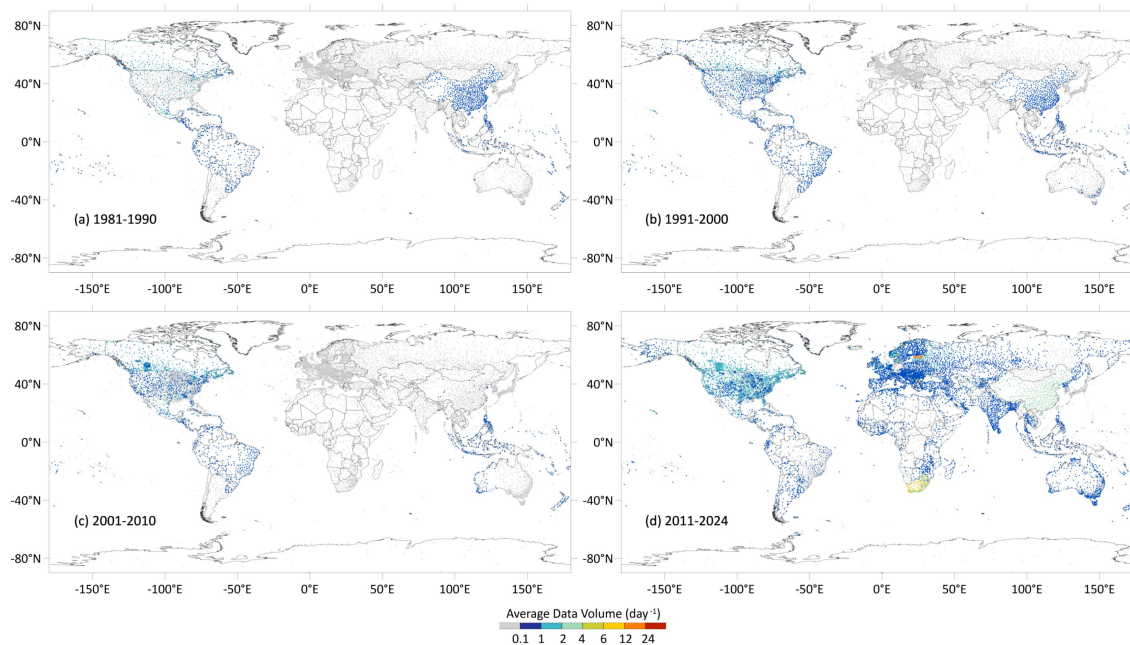
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Figure 4 Similar to Figure 2, but for Tmax-24h



203

Figure 5 Similar to Figure 1-S1, but for Tmin-12h



204

Figure 6 Similar to Figure 1-S1, but for Tmin-24h

205 3.2 GSOD in-situ daily data

206 The legacy GSOD dataset evaluated herein was accessed on 2025-03-01, reflecting
 207 the version that relied exclusively on the ISD as its input. The Global Surface
 208 Summary of the Day (GSOD; accessible at

<https://www.ncei.noaa.gov/access/metadata/landing-page/bin/iso?id=gov.noaa.ncdc:C00516>), maintained by the U.S. National Centers for Environmental Information (NCEI), generates daily meteorological summaries from over 9,000 global weather stations. Derived through systematic processing of the ISD, this product provides continuous records spanning 1929 to present, with post-1973 data exhibiting optimal completeness. Readers should be aware that with the retirement of the ISD, the legacy GSOD is slated for replacement by a new product, the Surface Summary of the Day (SSOD). Because the SSOD has not yet been officially released, a direct comparison with the true successor to GSOD is not currently possible. Furthermore, because GHCNv2 does not currently support our methodological requirements, real-time operationalization of GLBD-FED is temporarily paused.

4. Methods

Fig 7 illustrates the main procedures involved in the production of the Global Base Dataset-First Estimate Daily Data (GLBD-FED). The Integrated Surface Database (ISD) served as the data source, from which daily Tmax, Tave, and Tmin were derived using new methods after correcting for misrecorded data.

All daily data underwent rigorous quality control and were assigned specific quality and date boundary codes (detailed procedures are provided in Appendix, Section 11).

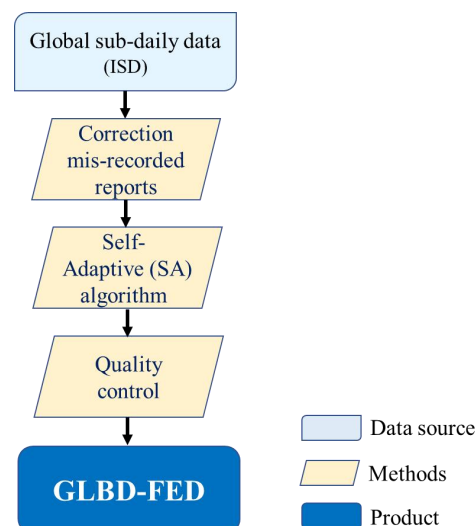


Figure 7 Main procedures of GLBD-FED production

4.1. Correction for Mis-recorded sub-daily Tmax/Tmin in data source

The quality of the data source is one of the most critical factors influencing the final data product. Notably, there were numerous misrecorded values for Tmax and Tmin over 12-hour and 24-hour periods in the global hourly data, particularly in South America.

Fig 8 presents a case study from San Antonio Oeste, Argentina (877840-99999), covering January 5 to January 12, 2023. The red and blue lines represent the derived upper limits of temperature from Tmax-12h and Tmax-24h, respectively, while the black circles indicate hourly Tave. Panel (a) displays results from the raw data, revealing that San Antonio Oeste tends to record Tmax-24h at 1200 UTC and Tmax-12h at 0000 and 1200 UTC. The blue lines (derived from Tmax-24h) are consistently lower than the hourly Tave, suggesting that these Tmax-24h values are likely erroneous due to incorrect recording.

The case study at San Antonio Oeste is highly representative of a systematic structural artifact within the legacy archive. A comprehensive global assessment of this specific mis-recording error from 1981 to 2024 reveals that it affected numerous stations worldwide, with concentrated occurrences in Western Europe, Russia, China, Canada, and South Africa. The primary driver of this anomaly is the misclassification of high-frequency sub-daily reports (e.g., 1-hour summaries) as 24-hour extremes. The spatial distribution and frequency of the stations affected are detailed in the Appendix (Figure A1). To ensure high-throughput scalability across the global network, these corrections were coded as automated conditional rules integrated directly into our processing pipeline. The specific programmatic procedures are as follows:

- Tmax-24h records at hh:00 were compared with the highest hourly Tave from the previous 12 hours (Tave-12hhhigh) and 24 hours (Tave-24hhhigh). If $Tmax-24h < Tave-24hhhigh - 0.5^{\circ}C$ and $Tmax-24h \geq Tave-12hhhigh$, it indicates that the Tmax-12h at hh:00 was incorrectly recorded as Tmax-24h and should be corrected.

- Tmax-12h records at hh:00 were similarly compared. If $Tmax-12h > Tave-12hhhigh + 0.5^{\circ}C$ and $Tmax-12h \leq Tave-24hhhigh$, it suggests that the Tmax-24h at hh:00 was inaccurately recorded as Tmax-12h and should also be corrected.

- The corrections for Tmin-24h and Tmin-12h followed a similar approach, but comparisons were made with the lowest hourly Tave from the previous 12 hours and 24 hours.

Fig 8 panel (b) displays the corrected results. The misrecorded Tmax-24h values were restored to Tmax-12h, and the derived hourly upper limit temperatures now align closely with the hourly Tave.

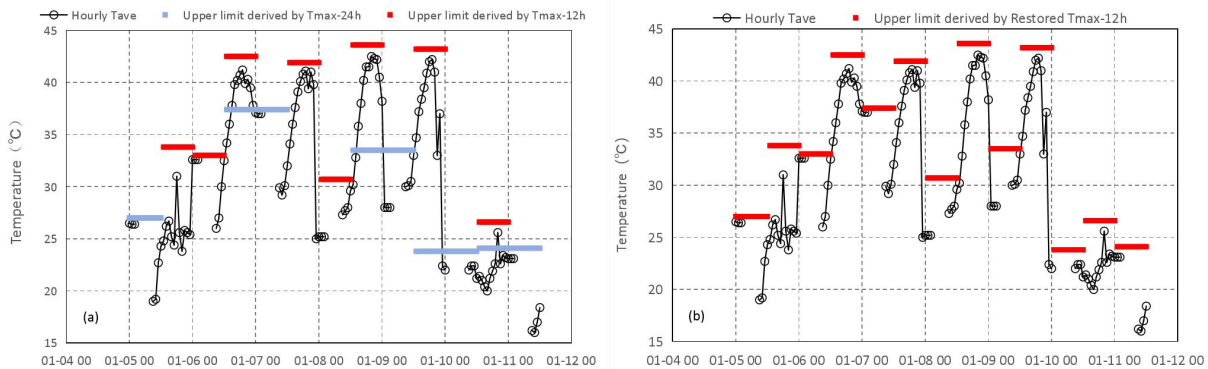


Figure 8 The discrete hourly observations and the derived upper limit of temperature deduced from Tmax-12h/24h during 5th Jan to 12th Jan 2023 at San Antonio Oeste, Argentina. Panels (a) and (b) represent the results from the raw and restored data, respectively. The misrecorded Tmax-24h (blue lines in panel a) are restored as Tmax-12h (red lines in panel b). (Note: The derived upper limits naturally plot slightly above the discrete hourly observations. This is not a visual offset, but reflects the physical reality that the continuous maximum temperature over a time window must be greater than or equal to any instantaneous hourly reading taken within that same period.)

4.2. New algorithms for in-situ daily temperature calculation

Aiming to obtain a set of global daily temperature data standing for the air temperature status under a prospective dateline, new algorithms were developed.

4.2.1. Time window management for daily data calculation

The standard 24-hour window for calculating daily variables (Tmax, Tave, and Tmin) is strictly defined as 0000 to 2400 UTC. However, to maximize global data retention, we implemented a sliding-window fallback mechanism. If sub-daily observations are insufficient to calculate a valid daily value within the standard UTC day, the 24-hour calculation window is systematically shifted by 1-hour increments (up to ± 12 hours) until a complete 24-hour data block is found. To ensure complete transparency and traceability, the specific start and end times utilized for each adjusted daily calculation are explicitly recorded in the dataset metadata.

4.2.2. Equidistant sampling for daily Tave calculation

The most analyzed variable in climate studies is mean annual or monthly temperature,

historically often defined using either fixed-hour observations or the simple average of daily maximum and minimum temperatures (Trewin, 2010). However, as discussed regarding spatiotemporal data volume (Section 3.1), explicit Tmax and Tmin reports do not always appear in pairs, exhibit lower geographic continuity than hourly observations, and are highly susceptible to Time of Observation Biases (TOB). This implies that deriving the daily Tave directly from the arithmetic mean of discrete, evenly distributed hourly records within the strict 0000–2400 UTC window is significantly more robust and globally uniform relative to deriving it from legacy extreme summaries. The daily average temperature is calculated as the following formula:

$$Tave_{daily}^1 = (Tave_{hh} + Tave_{hh+6hours} + Tave_{hh+12hours} + Tave_{hh+18hours}) / 4 \quad (1)$$

$hh = 0000, 0100, \dots, 0500 \text{ UTC}$

In this context, Tave represents the average temperature, with the subscript 'daily' indicating the daily aggregation and 'hh' denoting the specific starting hour. The value of hh is strictly one of the hours from 0000 to 0500 UTC (i.e., $hh \in \{0000, 0100, \dots, 0500\}$). In cases where higher-frequency sub-daily data provide multiple valid combinations, we apply a sequential priority rule: the algorithm selects the first available 6-hourly sequence by sequentially scanning the starting hours from 0000 UTC. To mitigate temporal sampling biases introduced by missing or irregular hourly observations, the daily average temperature (Tave) is only computed when there are at least 4 valid hourly observations relatively evenly distributed across the 24-hour period. Days failing to meet this data completeness threshold are recorded as missing.

4.2.3. Reaggregation for daily Tmax and Tmin calculation

Sub-daily Tmax and Tmin records were temporally decomposed into finer intervals, following the precipitation data processing protocols established by Yang et al. (2020). This approach facilitated probabilistic recombination into daily Tmax and Tmin values under a prospective dateline. Fig 9 illustrates a case of daily Tmin calculation using the reaggregation algorithm at False Pass, US (700638-99999) on October 11, 2024.

During the period from October 11 to October 12, False Pass recorded Tmin-12h at 1200 UTC and Tmin-24h at 0900 UTC. Initially, an implicitly deduced Tmin-15h at

0000 UTC on October 12 (-1.7°C , indicated by the dashed green arrow) was derived from the overlapping Tmin-24h at 0900 UTC (4.4°C , solid deep blue arrow) and Tmin-12h at 1200 UTC on October 12 (-1.7°C , solid light blue arrow). It is important to note that this temporal decomposition algorithm strictly requires perfect boundary alignment between the input records; if the boundaries do not perfectly align, the deduction becomes mathematically invalid and cannot be performed. This value was then combined with the Tmin-12h at 1200 UTC on October 11 (-1.1°C) to produce the daily Tmin value for October 11 (-1.7°C), which was the minimum recorded during that day.

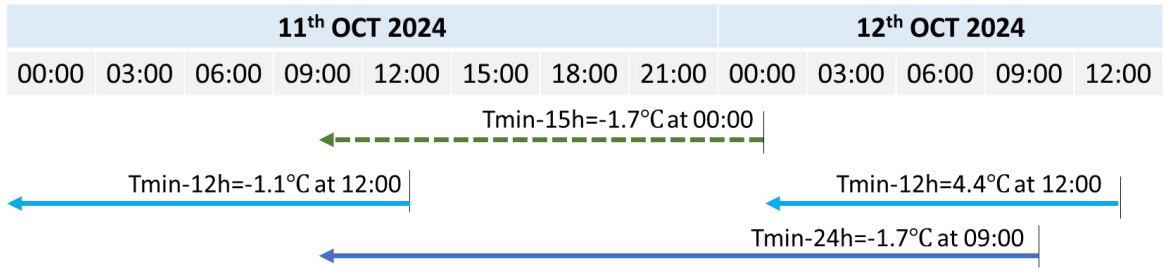


Figure 9 Application of the reaggregation algorithm for the daily Tmin value at False Pass, US (700638-99999) on October 11, 2024. The solid light and deep blue arrows represent the explicitly measured Tmin-12h and Tmin-24h records, respectively. The dashed green arrow represents the Tmin-15h implicitly deduced from these measured records.

In comparison to the original method (which used either two consecutive Tmax/Tmin records over 12 hours or one Tmax/Tmin record over 24 hours), the reaggregation algorithm demonstrates a significant improvement. Fig 10 illustrates the average data volume of global daily Tmax and Tmin obtained by the two methods in 2023. The volume of daily Tmax data increased by 64%, rising from approximately 3,400 to 5,600 entries per day, while the volume of daily Tmin data grew by 45%, increasing from 4,100 to 5,900 entries per day.

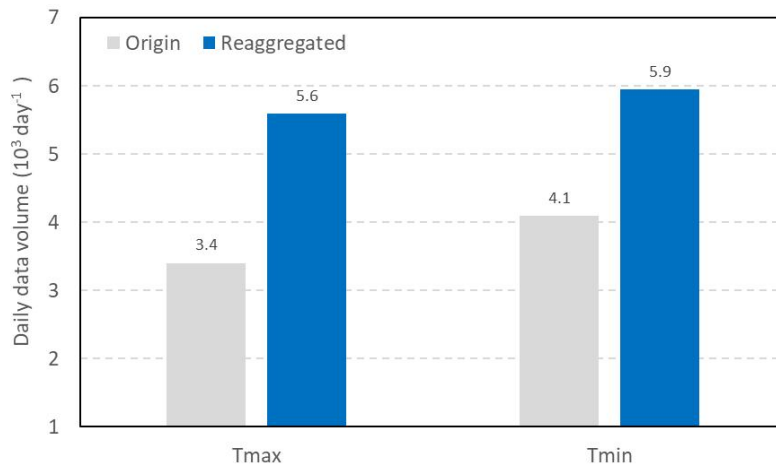
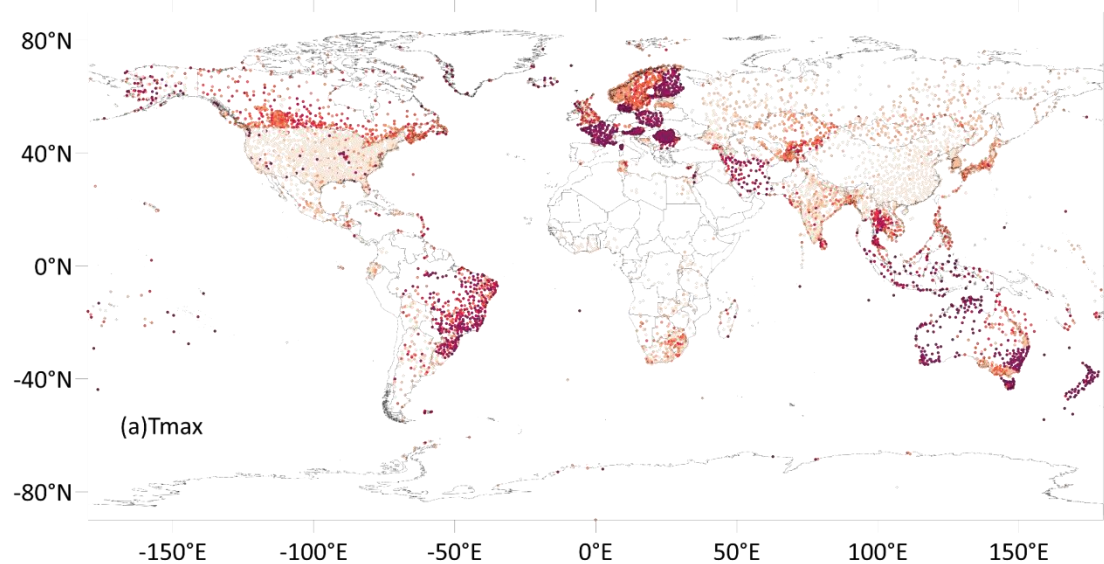


Figure 10 Average data volume of global daily Tmax and Tmin in 2023. The gray and blue bars

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represent the data calculated by origin and reaggregated methods, respectively.

337 Fig 11 illustrates the spatial distribution of the additional data recovered by our
338 temporal reconstruction algorithm. At the national level, several countries exhibit
339 particularly large increases. For instance, Canada, Brazil, Finland, Denmark, Poland,
340 Romania, Australia, New Zealand, Thailand, Indonesia, and the Philippines show
341 notable increases in retrieved daily Tmax records. Similarly, regions including Canada,
342 Brazil, Finland, Denmark, Poland, Romania, and Kazakhstan display significant
343 growth in daily Tmin. Overall, the recovery of Tmax exhibits a wider geographic
344 spread. This broader spatial rise is largely attributable to our algorithm successfully
345 resolving severe baseline scarcities caused by national reporting times clashing with
346 rigid UTC boundaries across networks in Australia and New Zealand. Notably, the
347 distribution reveals pronounced geopolitical border effects. These sharp spatial
348 contrasts at national boundaries directly reflect the differing sub-daily reporting
349 protocols and data exchange policies enforced by individual National Meteorological
350 and Hydrological Services (NMHSs). Because the temporal reconstruction algorithm
351 capitalizes on specific sub-daily reporting cadences, its enhancement effect naturally
352 maps onto these nation-specific administrative practices.



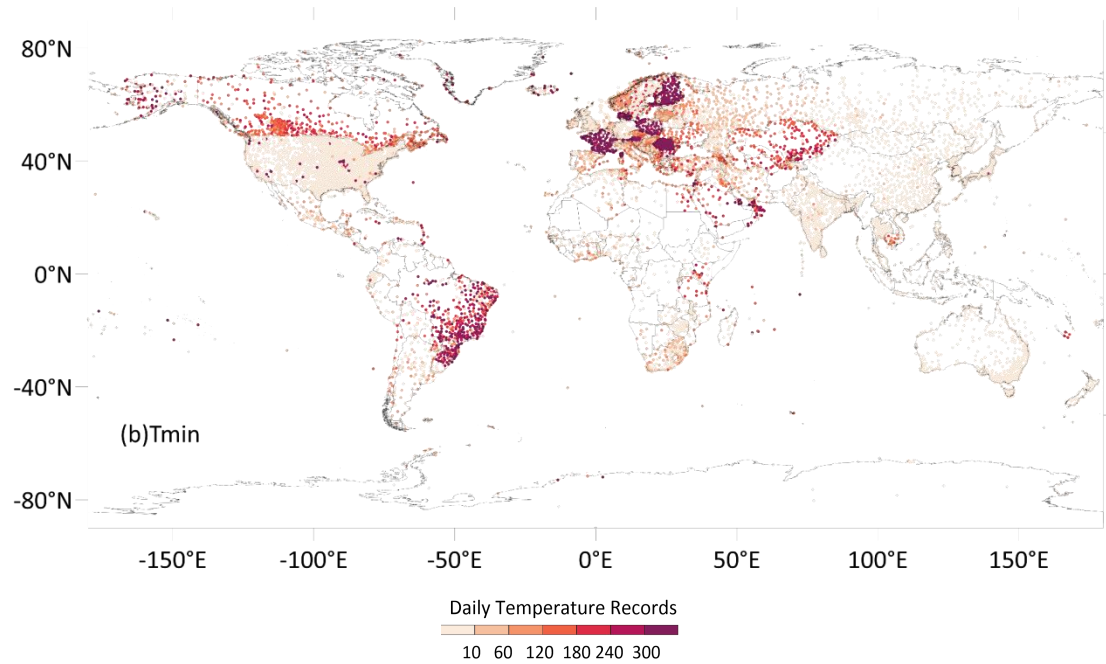


Figure 11 The spatial distribution of the increasing Tmax (panel a) and Tmin (panel b) data improved by the reaggregation algorithm in 2023. The color represents the increasing data volume

If the standard reaggregation algorithm fails or cannot be applied due to non-standard reporting times, a fallback strategy is implemented. First, the Tmax/Tmin-24h records that cover the greatest number of hours within the target day are used as the daily values, provided they overlap for at least 12 hours. Otherwise, the absolute maximum and minimum values derived from the discrete hourly temperature observations are employed as supplementary estimates for Tmax and Tmin, respectively, given that at least 21 valid hourly observations are available for that day. It is important to acknowledge that deriving daily extremes from discrete hourly temperature samples is subject to inherent sampling biases. Because true temperature extremes rarely occur exactly at the 'top' of the hour, estimates derived from quasi-instantaneous hourly observations tend to be slightly lower for Tmax and higher for Tmin than the true absolute values. Nevertheless, in the absence of explicit extreme reports, utilizing high-frequency hourly observations remains the only practical approach to maintain global temporal continuity.

4.3. Data quality controls

Given the massive volume of the global dataset and our objective to provide a rapid, scalable 'first estimate' baseline, the entire quality control procedure is fully automated. Quality controls were implemented for both the input hourly data and the

produced daily data. Global hourly temperature data quality tests include a spike value test, a stuck value test (which identifies prolonged sequences of the same value in the data series), and an inner consistency test (assessing the relationships between T_{max} , T_{ave} , and T_{min}). The daily data underwent these same tests, along with additional temporal and spatial consistency tests. Specifically, the temporal consistency test identifies values that deviate excessively from the station's long-term historical distribution, while the spatial consistency test compares a station's record with simultaneous observations from neighboring stations within a 100-km radius to detect localized anomalies. This specific suite of quality control tests was selected because it represents the established international standard, strictly adhering to WMO guidelines (WMO, 2011; 2018) and the comprehensive automated quality assurance protocols developed for major global datasets such as GHCN-Daily (Durre et al., 2010; Menne et al., 2012) and the ISD (Lott, 2004). Details of these tests for daily data are provided in the Appendix.

Throughout the processing pipeline, data quality results are evaluated via programmed conditional checks at each step and automatically flagged into three distinct categories: credible, suspicious, and erroneous. The quality test results at each stage are compiled into a final assessment of data quality levels for each individual data point (i.e., specifically for each individual daily record). A final quality level for a given observation is flagged as credible if there is no more than one suspicious test result and no erroneous test results across all applied tests. Conversely, a value is flagged as erroneous if more than one erroneous test result is found; otherwise, the final quality level is classified as suspicious.

5. Results and Discussion

In this section, we first present the temporal changes and spatial coverage of the global in-situ daily temperature data volume of GLBD-FED. This is followed by comparisons of data values between GLBD-FED and GSOD, along with further discussions.

5.1. Dataset Positioning and Current Status

It is important to emphasize that GLBD-FED is fundamentally designed as a

near-real-time, first-hand operational product rather than a retrospectively homogenized benchmark dataset. Its primary utility lies in rapid, temporally accurate evaluations of regional synoptic weather events and validating numerical weather prediction models. Caution should be exercised if applying it directly to long-term decadal climate trend detection without further statistical homogenization.

Furthermore, while the GLBD-FED processing framework was built for continuous near-real-time updates, its current operationalization is constrained by upstream data source transitions. Following the recent retirement of the ISD archive, we evaluated its successor, NOAA's GHCN-Hourly (GHCNh). Our assessment revealed that the meteorological elements contained in GHCNh have been significantly reduced, lacking the specific sub-daily extreme reports necessary to robustly support our daily Tmax and Tmin reconstruction methodology. Therefore, while real-time streaming is currently paused, the completed 1981–2024 GLBD-FED archive stands as a highly valuable, temporally aligned 44-year historical first-hand baseline for the global meteorological community.

5.2. In-situ Data Volume and Spatial Coverage

Fig 12 presents the spatial distribution (panels a1, a2, and a3) and temporal changes (panels b1, b2, and b3) of global daily Tmax, Tave, and Tmin data from 1981 to 2024. The colorful dots in panel a indicate the duration of daily data at each site, while the gray and black curves represent the daily data volume and the 15-point smoothing results, respectively.

Tmax, Tave, and Tmin exhibit very similar spatial distributions and temporal changes over the last four decades. Panels a1, a2, and a3 show approximately 17,000 sites with at least one year of daily Tmax data. Sites in China, Japan, and central Europe have extensive time series of daily Tmax (≥ 30 years, indicated by red dots). Western Europe and the US demonstrate a high spatial density of daily Tmax data.

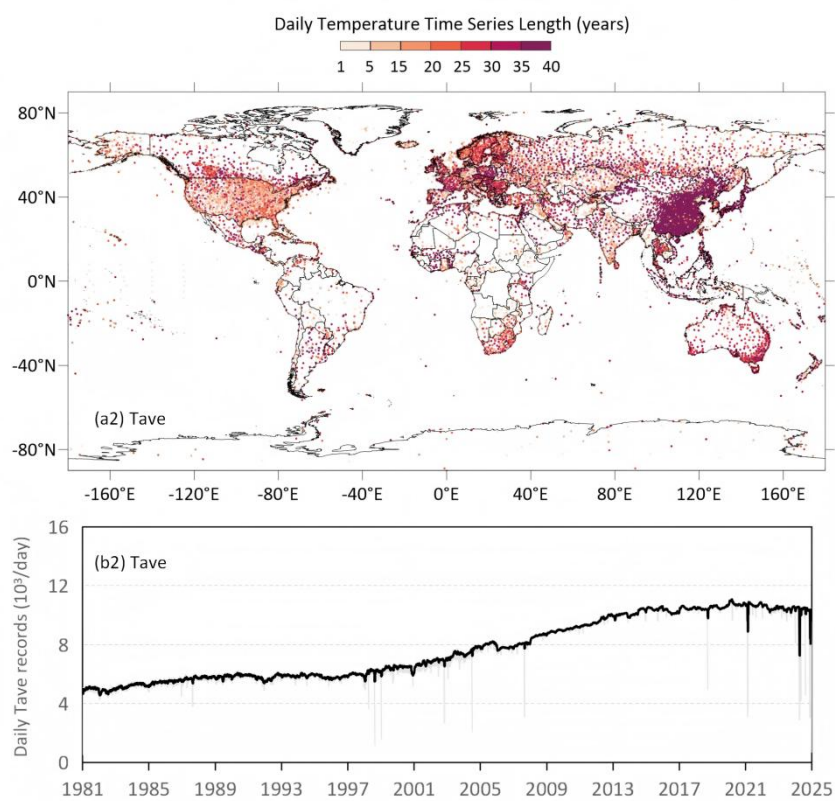
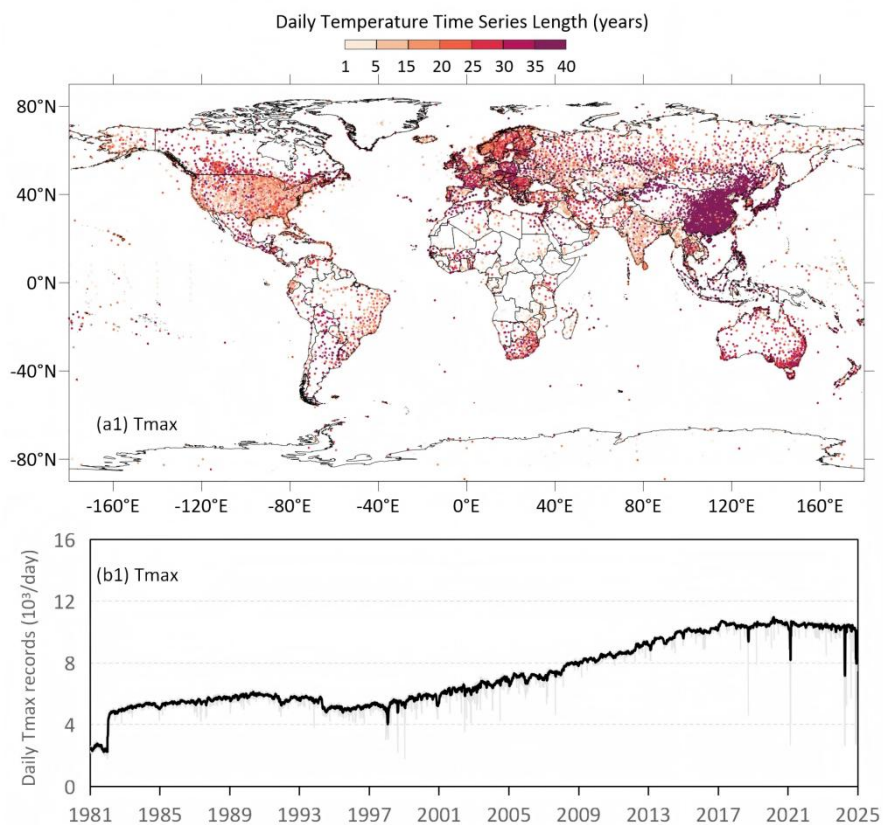
Notably, the apparent abundance of short time series in the U.S. (indicated by green and blue dots, < 20 years of records) is largely an artifact of the legacy ISD station identification system. Under this older system, minor station relocations or

administrative updates frequently resulted in the assignment of new station IDs, artificially fragmenting what are physically continuous observations. As recent advancements have shown, processing these records through the unified identifier framework of the newer GHCN_h dataset (e.g., merging closely located fragments under a single USW-prefix station ID) effectively resolves these artificial breaks and restores the long-term continuity of the time series. Conversely, countries like China tend to retain a consistent ID across relocations to preserve long-term continuity, which necessitates more rigorous homogenization attention in subsequent applications. Brazil also displays a high spatial density of data, with most stations beginning temperature observations in the last decade. Substantial sites with at least 20 years of daily temperature data can be found in southern Canada, coastal Australia, Russia, and southern Asia. Although the Antarctic and Arctic regions are among the most challenging for meteorological observation globally, dozens of sites have commenced measurements in the current century.

The global daily data volume for T_{max}, T_{ave}, and T_{min} has increased significantly over the last four decades. As shown in Figure 7 (panels b1, b2, and b3), the global daily data volume rose from approximately 3,000 entries per day in the 1980s to around 10,000 entries per day in recent years.

It should be noted that the occasional sharp drops observed in the time series are real features of the raw data stream, corresponding to specific periods of severe disruption in global data transmission caused by operational instabilities (e.g., telecommunication outages) within the WMO Global Telecommunication System (GTS) and national exchange networks.

Furthermore, the quasi-periodic fluctuations (dips) observed in the early T_{min} record (Fig. 7, panel b3) are artifacts of historical data scarcity in the underlying ISD archive. Our reaggregation algorithm is designed to primarily utilize explicitly reported extremes (T_{min}-12h/24h), relying on high-frequency hourly observations as a secondary fallback. During these earlier years, the raw archive occasionally lacked both explicit extreme reports and sufficient sub-daily observations. When neither data source meets the necessary computational thresholds, valid daily T_{min} values cannot be reconstructed, leading to temporary drops in the aggregated data volume.



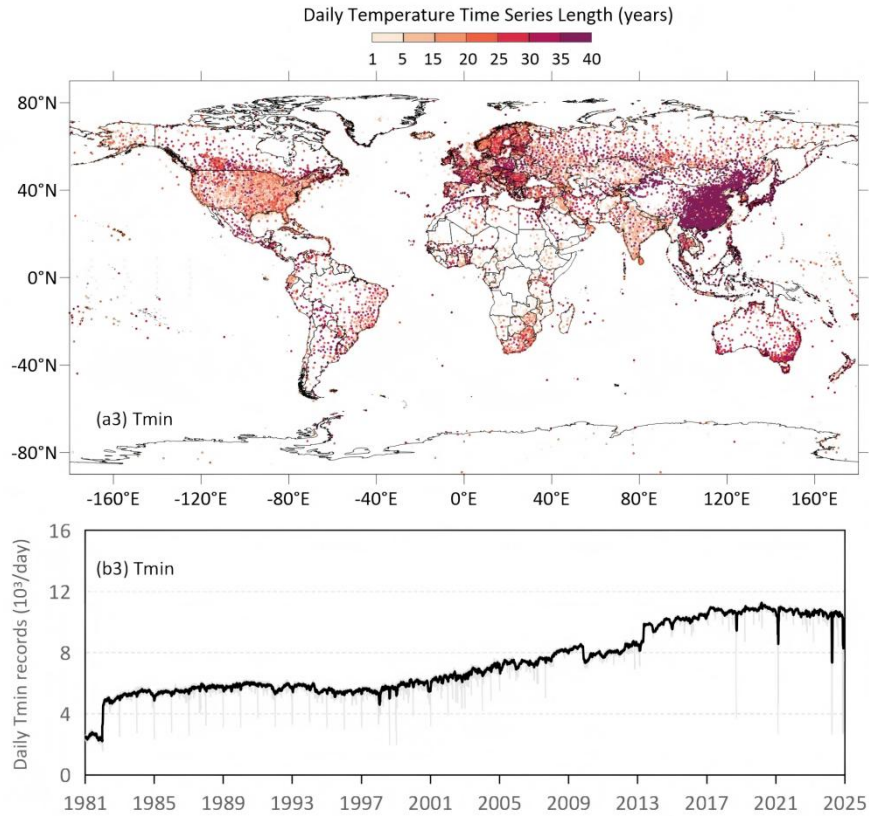


Figure 12 The spatial distribution (panels a1, a2, a3) and temporal changes (panels b1, b2, b3) of global daily temperature data during 1981–2024. Panels a1/b1 represent Tmax, a2/b2 represent Tave, a3/b3 represent Tmin. The colorful dots in panel a represent the length of daily data at sites (sequential 'rocket' color scale); the gray and black curves stand for the daily data volume and 15-point smoothing result.

5.3. Comparison of daily temperature with GSOD

Fig 13 presents the multi-decadal time series (1981–2024) of global daily temperature differences between GSOD and GLBD-FED. Calculated as the arithmetic mean across the strictly limited spatiotemporal intersection of the two datasets (i.e., only matched records present in both GSOD and GLBD-FED were included), the results demonstrate that GSOD exhibits a warmer daily Tmax (around +0.3°C), a colder Tmin (around -0.3°C), and nearly the same daily Tave (around +0.1°C) relative to GLBD-FED throughout the entire period. This means that rapid, near-real-time extreme weather assessments and operational synoptic evaluations relying on legacy datasets like GSOD would systematically misrepresent the magnitude of extreme events—overestimating warm extremes and underestimating cold extremes—compared to methodologically uniform evaluations based on GLBD-FED.

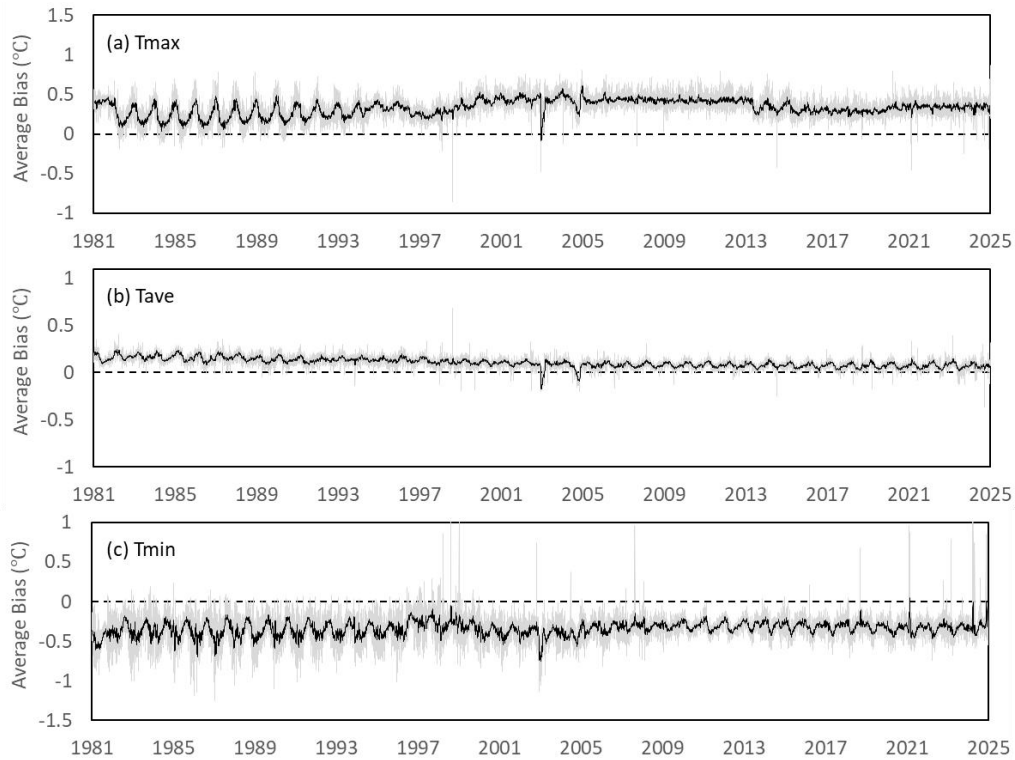
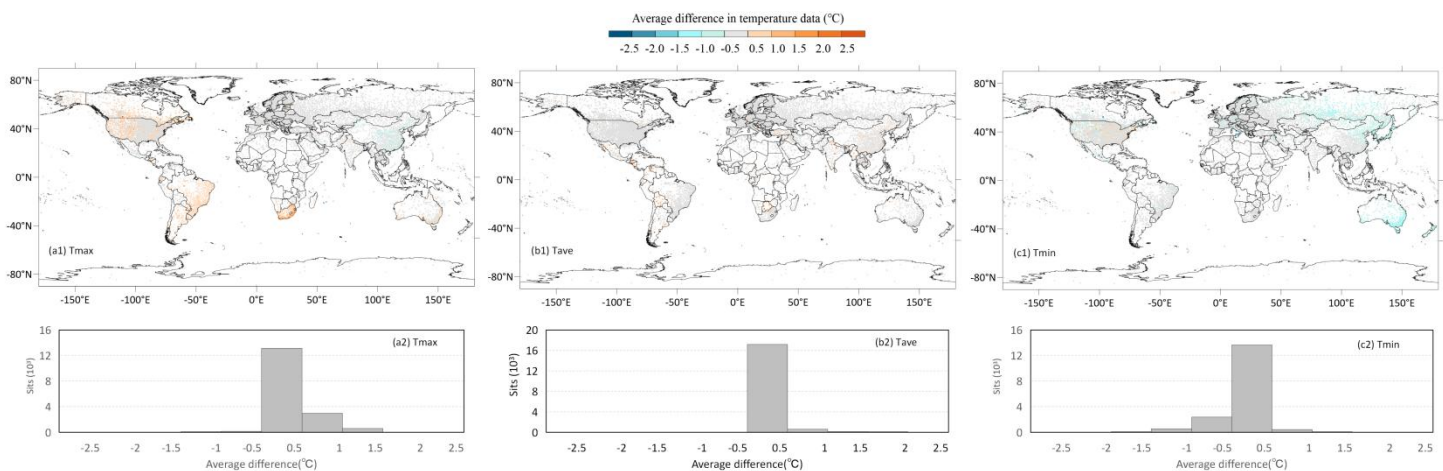


Figure 13 The daily time series of average difference in global in-situ daily temperature between GLBD-FED and GSOD (1981-2024) (GSOD minus GLBD-FED). The gray and black lines are the daily and 15 points-smoothing results, respectively. Panel (a), (b), and (c) stand for the results of Tmax, Tave and Tmin, respectively.

A detailed examination of the multi-decadal time series (Figure 8) reveals two notable temporal features: a pronounced seasonal cycle in the biases prior to 1990, and a long-term downward trend in the daily average temperature (Tave) bias (decreasing from approximately $+0.25^{\circ}\text{C}$ in the 1980s to roughly $+0.10^{\circ}\text{C}$ in the recent decade). Both features are intrinsically linked to historical data quality artifacts and algorithmic characteristics. In the early era, the seasonal cycle in the Tmax bias was largely driven by GSOD's processing of cold-season data from high-latitude regions. During boreal winters, raw observations in these environments were prone to anomalous positive spikes (unrealistic short-term jumps). GSOD's fallback extraction strategy misclassified these anomalous spikes as valid daily maximums, artificially inflating the winter Tmax and subsequently skewing the overall Tave upward. Simultaneously, the Tmin seasonality was driven by periodic variations in the volume of duplicated records within the GSOD archive. The long-term decreasing trend in the Tave bias reflects the progressive modernization of the global observing network. As automated weather stations and enhanced transmission protocols were widely deployed, the frequency of raw sensor noise dropped significantly. In contrast, GLBD-FED demonstrates higher stability throughout the 44-year period, as its rigorous temporal

501 consistency checks successfully filtered out these anomalous spikes even during the
 502 early, noisier era.

503 Fig 14 illustrates the spatiotemporal heterogeneity of daily temperature discrepancies
 504 between GSOD and GLBD-FED from 1981 to 2024. The sites in GSOD with higher
 505 daily Tmax ($\geq 0.5^{\circ}\text{C}$) relative to GLBD-FED are primarily located in Southern Africa,
 506 Brazil, Argentina, Canada, and the western United States (panel a1), accounting for
 507 21.4% of all sites (approximately 17,000) (panel a2).



508 Figure 14 Spatial distribution of the difference in in-situ daily temperature data between GLBD-FED
 509 and GSOD during 1981–2024 (GSOD minus GLBD-FED). Panels a1, b1, c1 show the difference in
 510 Tmax, Tave and Tmin at each site. Panels a2, b2, c2 represent the site number distribution with
 511 diversities.

512 Interestingly, GSOD shows lower Tmax values in China compared to GLBD-FED.
 513 This discrepancy arises from several types of reports labeled with the same ID in the
 514 ISD. The METAR and SYNOPTIC reports in China are independent and measured at
 515 different sites. METAR temperature reports tend to be lower than SYNOPTIC
 516 temperature reports since the latter are mostly recorded in urban areas. GSOD
 517 prioritized METAR reports as data sources, while GLBD-FED optimized Synoptic
 518 reports, leading to the observed negative bias in China.

519 Meanwhile, sites in GSOD with a lower daily Tmin ($\leq -0.5^{\circ}\text{C}$) compared to
 520 GLBD-FED are predominantly concentrated in Australia, Russia, and Northeast Asia
 521 (panel c1), accounting for 17.4% of all evaluated sites. For eastern Australia in
 522 particular, this regional bias is highly likely attributable to structural discrepancies

between national reporting practices and global transmission standards. Specifically, the standard national observation time for Tmin in eastern Australia is 2200 or 2300 UTC (depending on the season), whereas international SYNOPTIC reports sometimes stamp these observations at 0000 UTC. This temporal misalignment leads to Time of Observation Bias (TOB) causing double-counting or the shifting of minimum temperatures across the 0000 UTC boundary in the legacy GSOD dataset. Daily Tave from GSOD and GLBD-FED exhibits much greater consistency compared to Tmax and Tmin (panel b1), with 95.3% of sites maintaining differences within $\pm 0.5^{\circ}\text{C}$ (panel b2).

5.4. The causes for the discrepancy in daily temperature data

GLBD-FED and GSOD employ different algorithms to produce daily temperature data, resulting in discrepancies in data properties and biases between the two datasets, particularly for Tmax and Tmin. Table 1 outlines three main differences in the key processes of daily Tmax and Tmin data production between GLBD-FED and GSOD. First, GLBD-FED and GSOD have distinct definitions for daily Tmax and Tmin data. GLBD-FED considers records that could or almost represent the highest or lowest temperature in a 24-hour period as daily Tmax and Tmin, while GSOD selects the highest and lowest records within the 24 hours.

Second, although both GLBD-FED and GSOD utilize sub-daily data shared globally through the Global Telecommunication System (GTS) as their data source, they apply this data differently. GLBD-FED prefers SYNOPTIC reports (e.g., FM-12) over Meteorological Aerodrome (METAR, e.g., FM-15) and other reports when multiple types of sub-daily reports share the same ID, whereas GSOD does the opposite.

Third, GLBD-FED makes efforts to adjust the daily temperature boundary to 0000 UTC whenever possible, while GSOD retains the highest and lowest values as they appear within the 24-hour period.

Table 1 the comparison of the key processes in daily Tmax/Tmin calculation between GLBD-FED and GSOD

No	Key processes	GLBD-FED	GSOD
1	The definition of the daily Tmax/Tmin data	The records could/almost represent the highest/lowest	The highest/lowest records in the 24h.

		temperature over the 24h.	
2	The preference about the data source as there are several types of sub-daily reports labeled by the same ID	SYNOPTIC reports	Meteorological Aerodrome Reports
3	The date boundary for global daily data	Adjust the daily boundary of the daily temperature value to 00:00 UTC as much as possible.	/

5.4.1. The influence of the daily data definition

GLBD-FED and GSOD employ different methodologies for identifying daily Tmax and Tmin values. GLBD-FED calculates daily extremes by realigning sub-daily records to their physical 24-hour occurrence windows. In contrast, GSOD derives daily extremes using a hierarchical extraction approach—first selecting explicitly reported summaries, then falling back to discrete hourly observations—but attributes them based strictly on their UTC timestamps within the calendar day. This objective methodological difference is the primary driver of the systematic discrepancies observed between the two datasets. Consequently, in GSOD, a 24-hour extreme report that physically summarizes the previous day (e.g., a Tmax-24h reported at 0300 UTC) is erroneously treated as the daily extreme for the current day simply because of its timestamp. Furthermore, the theoretical sampling bias of hourly data provides an important context: since deriving extremes from hourly data inherently underestimates Tmax, the fact that GSOD still exhibits a significant global warm bias relative to GLBD-FED proves that GSOD's overestimation is overwhelmingly driven by its methodological characteristics (UTC-boundary double-counting and inclusion of anomalous data spikes), rather than sampling limitations.

Fig 15 displays the variations in air temperature at the SINPO station (470460) in North Korea from May 20 to May 23, 2014. The hollow black dots represent hourly average air temperatures recorded at fixed observation times, while the red and yellow dots indicate the Tmin-12h and Tmin-24h values recorded at 0000 and 2100 UTC, respectively. The light blue curve in panel (a) and the dark blue curve in panel (b) denote the lower limits derived by GSOD and GLBD-FED data, respectively.

In the GSOD dataset (panel a), the daily Tmin value for the SINPO station was recorded as 10.6°C for four consecutive days. This occurred because the Tmin-24h recorded at 0000 UTC on May 21, which was the lowest value during the period from

May 21 to May 22, was treated as the daily Tmin for both May 21 and May 22. As a result, the derived daily Tmin values for May 21 and May 23 are evidently underestimated and duplicated.

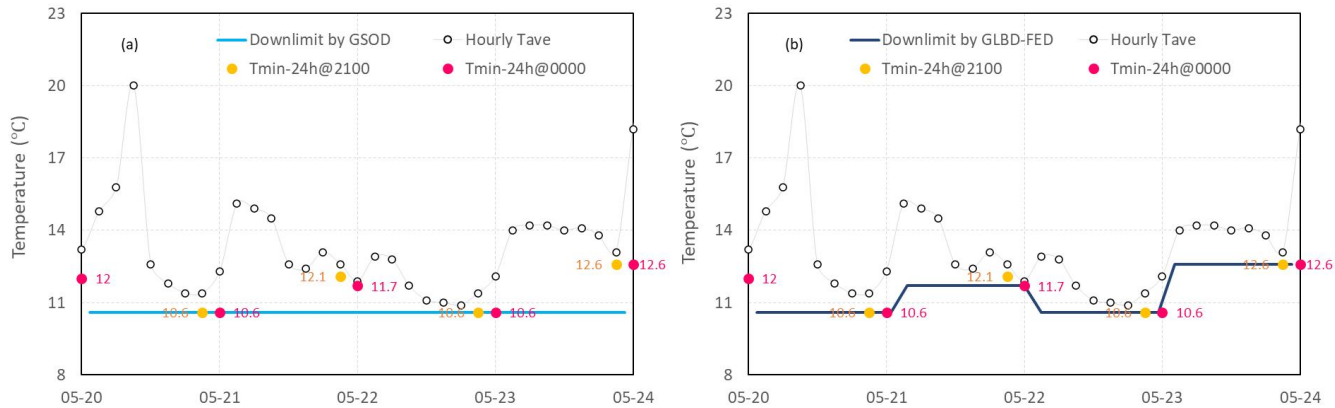
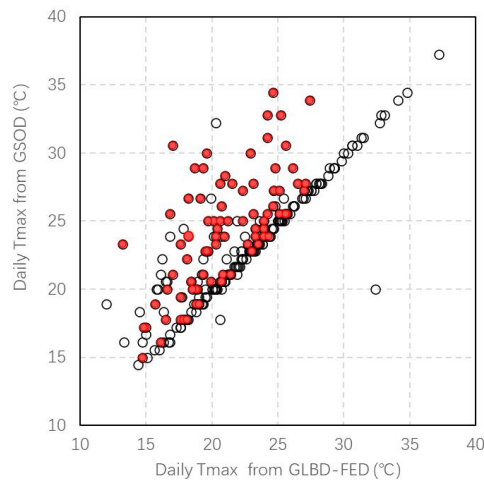


Figure 15 The lower-bound estimate of air temperature derived from the daily Tmin values based on GSOD (light blue, panel a) and GLBD-FED (dark blue, panel b) at the SINPO station (470460), North Korea, May 20–23, 2014.

Fig 16 presents a comparison of daily Tmax values at the Plettenberg Bay station in South Africa (689310-99999). The abscissa and ordinate represent the daily values from the GLBD-FED and GSOD datasets, respectively. The mean difference between the two sets of daily Tmax values is 1.3°C, indicating a mathematically higher overall daily Tmax in GSOD relative to GLBD-FED at this site. To thoroughly investigate the underlying cause, we categorized the data points into two groups based on the presence of consecutive identical values in GSOD (Fig 16). The black points (non-repeated values) predominantly fall along the 1:1 line, indicating zero or negligible difference on days without double-counting. However, approximately one-third of the Tmax daily values from GSOD this year were potential duplicated records (86 points, red dots), showing a positive mean difference of about 3.2°C relative to GLBD-FED. For context, an analysis of the GLBD-FED dataset for the same station and period reveals only 2 instances of consecutive identical values. Excluding the red points results in a roughly 70% reduction in the overall mean bias, decreasing it to 0.4°C. It is important to note that some repeated GSOD values (red points) lie exactly on the 1:1 line; these represent physically plausible weather conditions where the actual maximum temperature remained identical across consecutive days. Additionally, there are scattered points where GLBD-FED reports higher Tmax values than GSOD, typically when GLBD-FED's physical 24-hour window captures an extreme temperature event that is otherwise split across the UTC

603 boundary in GSOD.



604 Figure 16 A comparison of daily Tmax values at the Plettenberg Bay station (689310-99999), South
605 Africa, for the year 2024. The abscissa and ordinate represent the daily values from the GLBD-FED
606 and GSOD datasets, respectively. The red solid dots indicate the data that are repeated with their
607 previous ones (86 points), while the black hollow dots represent non-repeating values (172 points).

608 Furthermore, we investigated the extent of likely duplicated daily Tmax and Tmin
609 records in the GSOD dataset from 1981 to 2024 (identified as identical values derived
610 using the same calculation method on two consecutive days). Our analysis identified
611 approximately 18 million cases (averaging ~1,100 occurrences per day) of likely
612 duplicated Tmax records and 11 million cases (~670 occurrences per day) of likely
613 duplicated Tmin records. When compared against the GLBD-FED, these potential
614 duplicates introduced an average bias of 0.28°C and -0.85°C, respectively. This
615 indicates that the double-counting issue in legacy datasets systematically results in
616 artificially warmer maximum temperatures and colder minimum temperatures.

617 5.4.2. The influence of the choice of data source

618 There are several types of sub-daily reports labeled with the same ID in the ISD, and
619 these reports are not always measured at the same location. Both GLBD-FED and
620 GSOD treat these reports independently based on their types, but they have different
621 preferences. GLBD-FED prioritizes SYNOPTIC reports (e.g., FM-12) over METAR
622 reports (e.g., FM-15, FM-16) as data sources, whereas GSOD does the opposite. This
623 prioritization in GLBD-FED is driven by the goal of maximizing global dataset
624 uniformity. Synoptic observations are strictly coordinated by the WMO and adhere to

standardized global reporting protocols designed specifically for meteorological purposes. In contrast, while METAR (aviation) reports are abundant in specific regions such as the United States, their global distribution is highly uneven. Therefore, for a universal global dataset, prioritizing WMO-standardized SYNOPTIC reports ensures greater spatial and temporal homogeneity across different countries than relying on aviation-focused METAR data.

Table 2 presents a case study for Ulan Bator, Mongolia, where two specific types of sub-daily reports (FM-12 and FM-15) labeled under the identical ID 442920-99999 actually originate from different sites in the ISD. Figure 12 provides a targeted sensitivity analysis for Ulan Bator. Panel (a) displays the standard GLBD-FED daily Tmin calculated using our default preference for SYNOPTIC reports, while panel (b) shows the results when the algorithm is specifically forced to utilize only METAR reports from the same station ID. GSOD shows significantly lower daily Tmin values for Ulan Bator compared to GLBD-FED (SYNOPTIC reports), with an average bias of -4.5°C . The results in panel (b) nearly overlap with the GSOD data, indicating that the direct aggregation of airport-based METAR reports is a primary contributor to the significantly lower daily Tmin values recorded in GSOD for this identical station identifier.

Table 2 Ulan Bator, Mongolia (442920-99999) sub-daily reports

Reports Type	Longitude	Latitude	Altitude
SYNOPTIC Report (FM-12)	106.867°E	47.917°N	1306m
METAR Report (FM-15)	106.767°E	47.843°N	1330m

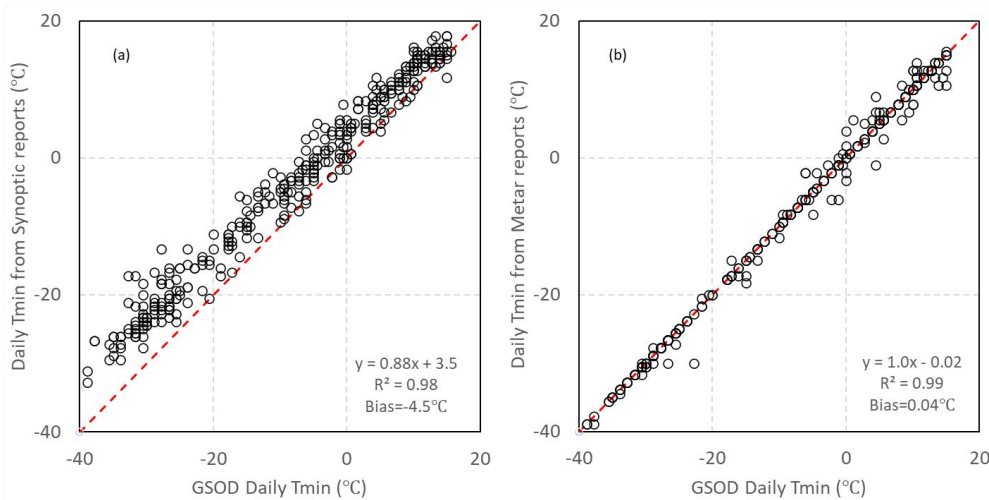


Figure 12 Comparison of daily Tmin at Ulan Bator, Mongolia (442920-99999) between GLBD-FED

and GSOD in 2024. Panel (a) compares the default GLBD-FED output (prioritizing SYNOPTIC reports, FM-12) with GSOD, while panel (b) shows a special comparison using GLBD-FED results derived exclusively from METAR reports (FM-15) from the same station ID.

5.4.3. The influence of the date boundary treatment

The systematic discrepancies identified in Fig. 8 originate from methodological divergences between GLBD-FED and GSOD, specifically in their temporal alignment frameworks. Fig 18(a) illustrates this mechanism through a case study at the Villa Reynolds station in Argentina (873280-99999; 31.96°S, 65.13°W), where GSOD exhibits a systematic Tmax warm bias of +1.6°C relative to GLBD-FED. This deviation occurs because GSOD attributes maximum temperatures to calendar days using the 1200 UTC Tmax-24h (0900 local time), in contrast to GLBD-FED's UTC-aligned approach, which reconstructs diurnal extremes by integrating next-day 1200 UTC observations with contemporaneous measurements. Temporal recalibration experiments (Fig 18(b)) effectively mitigate this discrepancy, achieving inter-dataset convergence at 0.1°C with improved correlation ($R^2=0.99$ compared to the original $R^2=0.87$).

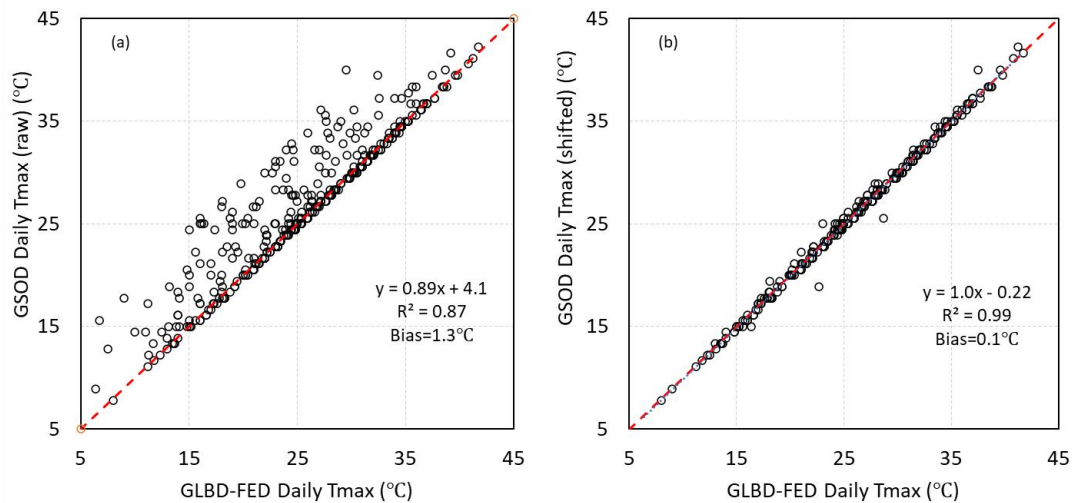


Figure 18 The comparison of daily Tmax at Villa Reynolds Argentina (873280-99999) between GLBD-FED and GSOD in 2022. Panel a is the comparison between GLBD-FED and original GSOD data, while panel b is the comparison between GLBD-FED and date shifted (forward moving) GSOD data.

6. Conclusion

This study develops GLBD-FED, a global in situ daily temperature dataset encompassing maximum (Tmax), average (Tave), and minimum (Tmin) temperatures, constructed using quasi-real-time sub-daily observations from ISD. The main results

of this study show that:

1. To produce a global daily dataset representing maximum, minimum, and average temperatures over a rigorously defined 24-hour period (i.e., the standard 0000–2400 UTC day), we developed a new algorithm that decomposes sub-daily Tmax and Tmin records into finer intervals and then reaggregates them into daily extremes under physical 24-hour calculation windows. Compared to the conventional calculation method (which relied on either two consecutive Tmax/Tmin records over 12 hours or one Tmax/Tmin record over 24 hours), the new algorithm significantly increased the data counts of valid daily Tmax and Tmin records by 64% and 45%, respectively. A correction for misrecorded Tmax and Tmin records was also implemented.

2. GLBD-FED includes Tmax, Tave, and Tmin data from approximately 17,000 global sites covering the period from 1981 to 2024. The daily temperature volume of GLBD-FED increased from 3,000 records per day in the 1980s to 10,000 records per day in the 2020s. America and Asia show high spatial densities of daily temperature data, especially in recent years. In comparison to GSOD, Tmax and Tmin from GLBD-FED exhibit less extreme daily values, with slightly lower daily Tmax (approximately -0.30°C) and higher daily Tmin (approximately $+0.30^{\circ}\text{C}$), resulting in nearly the same daily Tave (around $+0.10^{\circ}\text{C}$). These differences are primarily attributed to multiple sources, including the diversity in daily data definitions, the choice of data sources, and date boundary treatment methods.

7. Data availability

The global in-situ temperature daily data are publicly available at <https://zenodo.org/records/17895292> (Yang et al., 2025).

8. Author contributions

SY designed and carried out the study, performed the analyses, and drafted the manuscript. PMZ and XBZ contributed to the development of the manuscript framework, provided scientific guidance throughout the research process, and critically reviewed and revised the manuscript. HJ and FY contributed to the preprocessing of the data source and the automated data production processes,

respectively. ZJZ reviewed and edited the manuscript.

9. Competing interests

The contact authors have declared that none of the authors has any competing interests.

10. Acknowledgments

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The authors also would like to express their sincere gratitude to the three reviewers for their highly constructive comments and rigorous evaluation, which significantly improved the quality and clarity of this manuscript. We also extend our deep appreciation to the Topic Editor and the Chief Editor of Earth System Science Data for their efficient handling of the manuscript and their valuable guidance throughout the peer-review process.

11. Appendix

11.1. The spatial distribution of percentage of mis-recording Tmax and Tmin data

Figure A1 illustrates the spatial distribution of the percentage of mis-recorded Tmax and Tmin data from 1981 to 2024 (mentioned in Section 4.1). These anomalies are distributed across multiple regions globally, with notable concentrations in Western

Europe and South Africa. The primary cause of these errors is the frequent misclassification of high-frequency Tmax reports (e.g., 1-hour interval summaries) as 24-hour extreme values, an artifact that is particularly prevalent across Europe.

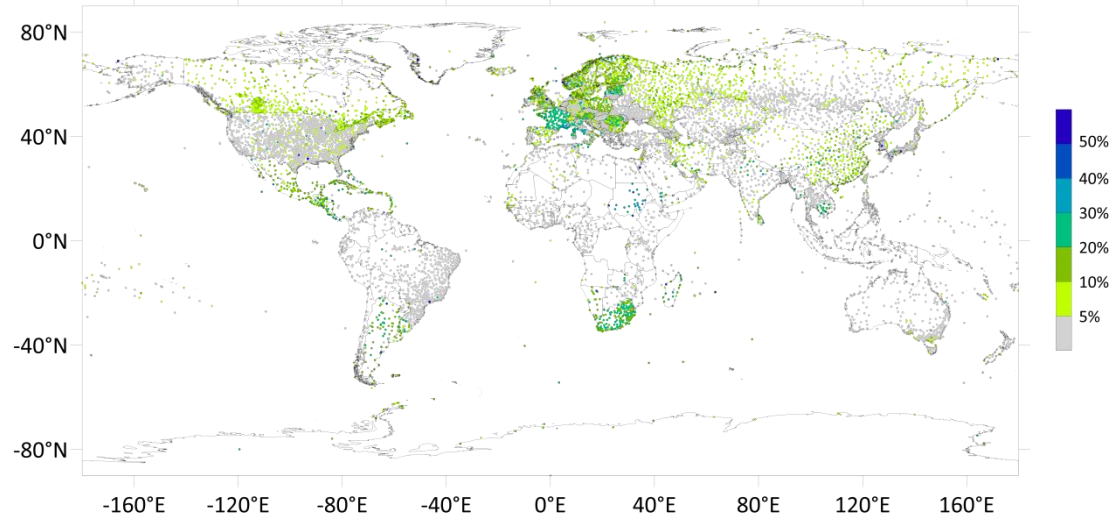


Figure A1 The spatial distribution of percentage of mis-recording Tmax and Tmin data from 1981 to 2024.

11.2. Formulae used for Daily Data Quality Data Quality Tests

Spike test

This test identifies anomalous daily temperature values that exceed predefined absolute physical boundaries or historical climatological limits for a specific station. Where the subscript i stands for the measurement on the i -th day; $x_{upperlimit}$ / $x_{lowerlimit}$ is the upper/lower threshold value for the record and is the smaller/higher value between $\max(\bar{x}) + 5^{\circ}\text{C}$ and 80°C , where $\bar{x}$ represents the subset of the historic measurements in the month (Jan, Feb,...Dec) which removes the smallest and largest 1% of values.

$$qc = \begin{cases} \text{credible}, & x_{lowerlimit} \leq x_i \leq x_{upperlimit} \\ \text{wrong}, & x_i \geq x_{upperlimit} \text{ or } x_i \leq x_{lowerlimit} \end{cases}$$

$$x_{upperlimit} = \min[\max(\bar{x}) + 5^{\circ}\text{C}, 80^{\circ}\text{C}]$$

$$x_{lowerlimit} = \max[\min(\bar{x}) - 5^{\circ}\text{C}, -80^{\circ}\text{C}]$$

Inner consistency test

This test verifies the basic logical relationship among the three daily temperature variables. It ensures that the daily maximum temperature is greater than or equal to

741 the daily average temperature, which in turn must be greater than or equal to the daily
 742 minimum temperature. Where T_{max} , T_{ave} , T_{min} stand for the daily T_{max} , T_{ave} ,
 743 T_{min} data.

$$qc = \begin{cases} credible, T_{max} \geq T_{ave} \geq T_{min}, \\ wrong, T_{max} \leq T_{ave} \leq T_{min} \\ doubtful, else \end{cases}$$

744 **Temporal consistency test**

745 This test detects daily records that deviate excessively from the historical statistical
 746 distribution of a given station, evaluated using the median and standard deviation.
 747 Where the μ and σ are the median value and standard deviation of $\bar{x}$, respectively.

$$qc = \begin{cases} credible, x_i \leq \mu + 2.5\sigma \\ wrong, x_i > \mu + 5\sigma \\ doubtful, \mu + 2.5\sigma < x_i \leq \mu + 5\sigma \end{cases}$$

$$\mu = \text{median}(\bar{x})$$

$$\sigma = \text{std}(\bar{x})$$

748 **Spatial consistency test**

749 This test identifies records that exhibit significant discrepancies when compared to
 750 concurrent observations from neighboring stations within the same region. Where the
 751 subscript j stands for the j -th neighbouring site (within the 100 km radius around the
 752 candidate site) and m is the total number of neighbouring sites.

$$qc = \begin{cases} credible, \delta_i \leq \min [\bar{\delta} + 2.5\nabla\delta, 10^\circ\text{C}] \\ doubtful, \min [\bar{\delta} + 2.5\nabla\delta, 10^\circ\text{C}] < \delta_i < \min [\bar{\delta} + 3.5\nabla\delta, 15^\circ\text{C}] \\ wrong, \delta_i \geq \min [\bar{\delta} + 3.5\nabla\delta, 15^\circ\text{C}] \end{cases}$$

$$\delta_i = \sqrt{\frac{\sum_{j=1}^m (x_j - x_i)^2}{m}}$$

$$\bar{\delta} = \text{median}(\delta_i)$$

$$\nabla\delta = \text{std}(\delta_i)$$

753 **Repeat test**

754 This test identifies instances where identical temperature values are recorded
 755 consecutively over a specified number of days. Where T_s and T_a are the standard
 756 deviation and smoothing average of 3 days measurements, respectively.

$$qc = \begin{cases} wrong, \sigma = 0 \\ credible, else \end{cases}$$

$$\sigma = \text{std}(x_{i-1}, x_i, x_{i+1})$$

$$\mu = \text{mean}(x_{i-1}, x_i, x_{i+1})$$

where σ and μ are the standard deviation and smoothing average of 3 days measurements, respectively.

12. Reference

- Adler, R. F., Huffman, G. J., Chang, A., Ferraro, R., Xie, P.-P., Janowiak, J., Rudolf, B., Schneider, U., Curtis, S., Bolvin, D., et al.: The Version-2 Global Precipitation Climatology Project (GPCP) monthly precipitation analysis (1979–Present), *J. Hydrometeorol.*, 4, 1147–1167, [https://doi.org/10.1175/1525-7541\(2003\)004<1147:TVGPCP>2.0.CO;2](https://doi.org/10.1175/1525-7541(2003)004<1147:TVGPCP>2.0.CO;2), 2003.
- Alexander, L. V., Zhang, X., Peterson, T. C., Caesar, J., Gleason, B., Klein Tank, A. M. G., Haylock, M., Collins, D., Trewin, B., Rahimzadeh, F., et al.: Global observed changes in daily climate extremes of temperature and precipitation, *J. Geophys. Res.-Atmos.*, 111, D05109, <https://doi.org/10.1029/2005JD006290>, 2006.
- Beck, H. E., van Dijk, A. I. J. M., Levizzani, V., Schellekens, J., Miralles, D. G., Martens, B., and de Roo, A.: MSWEP: 3-hourly 0.25° global gridded precipitation (1979–2015) by merging gauge, satellite, and reanalysis data, *Hydrol. Earth Syst. Sci.*, 21, 589–615, <https://doi.org/10.5194/hess-21-589-2017>, 2017.
- Beck, H. E., Wood, E. F., Pan, M., Fisher, C. K., Miralles, D. G., van Dijk, A. I. J. M., McVicar, T. R., and Adler, R. F.: MSWEP V2 global 3-hourly 0.1° precipitation: Methodology and quantitative assessment, *Bull. Am. Meteorol. Soc.*, 100, 473–500, <https://doi.org/10.1175/BAMS-D-17-0138.1>, 2019.
- Becker, A., Finger, P., Meyer-Christoffer, A., Rudolf, B., Schamm, K., Schneider, U., and Ziese, M.: A description of the global land-surface precipitation data products of the GPCC with sample applications including centennial trend analysis from 1901–present, *Earth Syst. Sci. Data*, 5, 71–99, <https://doi.org/10.5194/essd-5-71-2013>, 2013.
- Blenkinsop, S., Fowler, H. J., Barbero, R., Chan, S. C., Guerreiro, S. B., Kendon, E., Lenderink, G., Lewis, E., Li, X.-F., Westra, S., et al.: The INTENSE project: Using observations and models to understand the past, present and future of sub-daily rainfall extremes, *Adv. Sci. Res.*, 15, 117–126, <https://doi.org/10.5194/asr-15-117-2018>, 2018.
- Caesar, J., Alexander, L., and Vose, R.: Large-scale changes in observed daily

maximum and minimum temperatures: Creation and analysis of a new gridded data set, *J. Geophys. Res.-Atmos.*, 111, D05101, <https://doi.org/10.1029/2005JD006280>, 2006.

Daly, C.: Variable influence of terrain on precipitation patterns: Delineation and use of effective terrain height in PRISM, PRISM Group, Oregon State University, Corvallis, OR, USA, 1–7, 2002.

Dietzsch, F., Andersson, A., Ziese, M., Schröder, M., Raykova, K., Schamm, K., and Becker, A.: A global ETCCDI-based precipitation climatology from satellite and rain gauge measurements, *Climate*, 5, 9, <https://doi.org/10.3390/cli5010009>, 2017.

Dunn, R. J. H., Alexander, L. V., Donat, M. G., Zhang, X., Bador, M., Herold, N., Lippmann, T., Allan, R., Aguilar, E., Barry, A. A., et al.: Development of an updated global land in situ-based data set of temperature and precipitation extremes: HadEX3, *J. Geophys. Res.-Atmos.*, 125, e2019JD032263, <https://doi.org/10.1029/2019JD032263>, 2020.

Donat, M. G., Alexander, L. V., Yang, H., Durre, I., Vose, R., Dunn, R. J. H., Willett, K. M., Aguilar, E., Brunet, M., Caesar, J., et al.: Updated analyses of temperature and precipitation extreme indices since the beginning of the twentieth century: The HadEX2 dataset, *J. Geophys. Res.-Atmos.*, 118, 2098–2118, <https://doi.org/10.1002/jgrd.50150>, 2013.

Durre, I., Menne, M. J., Gleason, C. E., Houston, T. G., and Vose, R. S.: Comprehensive automated quality assurance of daily surface observations, *J. Appl. Meteorol. Climatol.*, 49, 1615–1633, <https://doi.org/10.1175/2010JAMC2375.1>, 2010.

Haiden, T., Janousek, M., Bidlot, J., Buizza, R., Ferranti, L., Prates, F., and Vitart, F.: Evaluation of ECMWF forecasts, including the 2018 upgrade, ECMWF Technical Memorandum No. 831, ECMWF, Reading, UK, <https://doi.org/10.21957/b8g71b69>, 2018.

Harris, I., Jones, P. D., Osborn, T. J., and Lister, D. H.: Updated high-resolution grids of monthly climatic observations – the CRU TS3.10 Dataset, *Int. J. Climatol.*, 34, 623–642, <https://doi.org/10.1002/joc.3711>, 2014.

Harris, I., Osborn, T. J., Jones, P., and Lister, D.: Version 4 of the CRU TS monthly high-resolution gridded multivariate climate dataset, *Sci. Data*, 7, 109, <https://doi.org/10.1038/s41597-020-0453-3>, 2020.

Hulme, M.: An intercomparison of model and observed global precipitation climatologies, *Geophys. Res. Lett.*, 18, 1715–1718,

822 <https://doi.org/10.1029/91GL01506>, 1991.

823 Hulme, M.: A 1951–80 global land precipitation climatology for the evaluation of
824 general circulation models, *Clim. Dyn.*, 7, 57–72,
825 <https://doi.org/10.1007/BF00209609>, 1992.

826 Janowiak, J. E., Gruber, A., Kondragunta, C. R., Livezey, R. E., and Huffman, G. J.: A
827 comparison of the NCEP–NCAR reanalysis precipitation and the GPCP rain
828 gauge–satellite combined dataset with observational error considerations, *J. Clim.*, 11,
829 2960–2979, [https://doi.org/10.1175/1520-0442\(1998\)011<2960:ACOTNN>2.0.CO;2](https://doi.org/10.1175/1520-0442(1998)011<2960:ACOTNN>2.0.CO;2),
830 1998.

831 Jones, P. D. and Hulme, M.: Calculating regional climatic time series for temperature
832 and precipitation: Methods and illustrations, *Int. J. Climatol.*, 16, 361–377,
833 [https://doi.org/10.1002/\(SICI\)1097-0088\(199604\)16:4<361::AID-JOC53>3.0.CO;2-F](https://doi.org/10.1002/(SICI)1097-0088(199604)16:4<361::AID-JOC53>3.0.CO;2-F),
834 1996.

835 Jones, P. D. and Moberg, A.: Hemispheric and large-scale surface air temperature
836 variations: An extensive revision and an update to 2001, *J. Clim.*, 16, 206–223,
837 [https://doi.org/10.1175/1520-0442\(2003\)016<0206:HALSSA>2.0.CO;2](https://doi.org/10.1175/1520-0442(2003)016<0206:HALSSA>2.0.CO;2), 2003.

838 Kamiguchi, K., Arakawa, O., Kitoh, A., Yatagai, A., Hamada, A., and Yasutomi, N.:
839 Development of APHRO_JP, the first Japanese high-resolution daily precipitation
840 product for more than 100 years, *Hydrol. Res. Lett.*, 4, 60–64,
841 <https://doi.org/10.3178/hrl.4.60>, 2010.

842 Karl, T. R., Williams Jr, C. N., Young, P. J., and Wendland, W. M.: A model to
843 estimate the time of observation bias associated with monthly mean maximum,
844 minimum and mean temperatures for the United States, *J. Clim. Appl. Meteorol.*, 25,
845 145–160, [https://doi.org/10.1175/1520-0450\(1986\)025<0145:AMTETT>2.0.CO;2](https://doi.org/10.1175/1520-0450(1986)025<0145:AMTETT>2.0.CO;2),
846 1986.

847 Klein Tank, A. M. G., Wijngaard, J. B., Können, G. P., Böhm, R., Demarée, G., Goma,
848 A., Mileta, M., Pashiardis, S., Hejkrlik, L., Kern-Hansen, C., et al.: Daily dataset of
849 20th-century surface air temperature and precipitation series for the European Climate
850 Assessment, *Int. J. Climatol.*, 22, 1441–1453, <https://doi.org/10.1002/joc.773>, 2002.

851 Lott, N.: The quality control of the Integrated Surface Database (ISD), in: Preprints,
852 84th AMS Annual Meeting, American Meteorological Society, Seattle, WA, USA, 1–7,
853 2004.

854 Menne, M. J., Durre, I., Vose, R. S., Gleason, B. E., and Houston, T. G.: An overview
855 of the Global Historical Climatology Network-Daily Database, *J. Atmos. Ocean.*

Technol., 29, 897–910, <https://doi.org/10.1175/JTECH-D-11-00103.1>, 2012.

Menne, M. J., Noone, S., Casey, N., Dunn, R., McNeill, S., Kantor, D., Thorne, P., Orcutt, K., Cunningham, S., and Risavi, N.: The Global Historical Climatology Network hourly (GHCNh) dataset, in: 105th AMS Annual Meeting, American Meteorological Society, New Orleans, LA, USA, 2025.

Mitchell, T. D. and Jones, P. D.: An improved method of constructing a database of monthly climate observations and associated high-resolution grids, *Int. J. Climatol.*, 25, 693–712, <https://doi.org/10.1002/joc.1181>, 2005.

National Centers for Environmental Information (NCEI): Global Surface Summary of the Day (GSOD), NOAA [data set], 2017.

New, M., Hulme, M., and Jones, P.: Representing twentieth-century space–time climate variability. Part I: Development of a 1961–90 mean monthly terrestrial climatology, *J. Clim.*, 12, 829–856, [https://doi.org/10.1175/1520-0442\(1999\)012<0829:RTCSTC>2.0.CO;2](https://doi.org/10.1175/1520-0442(1999)012<0829:RTCSTC>2.0.CO;2), 1999.

New, M., Hulme, M., and Jones, P.: Representing twentieth-century space-time climate variability. Part II: Development of 1901–96 monthly grids of terrestrial surface climate, *J. Clim.*, 13, 2217–2238, [https://doi.org/10.1175/1520-0442\(2000\)013<2217:RTCSTC>2.0.CO;2](https://doi.org/10.1175/1520-0442(2000)013<2217:RTCSTC>2.0.CO;2), 2000.

Rohde, R., Muller, R. A., Jacobsen, R., Muller, E., Perlmutter, S., Rosenfeld, A., Wurtele, J., Groom, D., and Wickham, C.: A new estimate of the average Earth surface land temperature spanning 1753 to 2011, *Geoinfor. Geostat. An Overview*, 1, 1000101, <https://doi.org/10.4172/2327-4581.1000101>, 2013.

Rudolf, B.: Management and analysis of precipitation data on a routine basis, in: *Proceedings of the International Symposium on Precipitation and Evaporation*, Bratislava, Slovakia, 20–24 September 1993, 69–76, 1993.

Schamm, K., Ziese, M., Becker, A., Finger, P., Meyer-Christoffer, A., Schneider, U., Schröder, M., and Stender, P.: Global gridded precipitation over land: A description of the new GPCC First Guess Daily product, *Earth Syst. Sci. Data*, 6, 49–60, <https://doi.org/10.5194/essd-6-49-2014>, 2014.

Schneider, U.: The GPCC quality control system for gauge measured precipitation data, WCP-115, WMO/TD-No. 718, World Meteorological Organization (WMO), Geneva, Switzerland, 1993.

Schneider, U., Becker, A., Ziese, M., and Rudolf, B.: Global precipitation analysis products of the GPCC, Global Precipitation Climatology Centre (GPCC), DWD,

890 Offenbach a.M., Germany, 1–14, 2018.

891 Smith, A., Lott, N., and Vose, R.: The Integrated Surface Database: Recent
 892 developments and partnerships, *Bull. Am. Meteorol. Soc.*, 92, 704–708,
 893 <https://doi.org/10.1175/2011BAMS3015.1>, 2011.

894 Trewin, B.: Exposure, instrumentation, and observing practice effects on land
 895 temperature measurements, *Wiley Interdiscip. Rev.: Clim. Change*, 1, 490–506,
 896 <https://doi.org/10.1002/wcc.46>, 2010.

897 Westra, S., Fowler, H. J., Evans, J. P., Alexander, L. V., Berg, P., Johnson, F., Kendon,
 898 E. J., Lenderink, G., and Roberts, N. M.: Future changes to the intensity and
 899 frequency of short-duration extreme rainfall, *Rev. Geophys.*, 52, 522–555,
 900 <https://doi.org/10.1002/2014RG000464>, 2014.

901 World Meteorological Organization (WMO): Guide to Climatological Practices,
 902 WMO-No. 100, World Meteorological Organization, Geneva, Switzerland, 2011.

903 World Meteorological Organization (WMO): Manual on Codes – International Codes,
 904 Volume I.1, Part A – Alphanumeric Codes, WMO-No. 306, World Meteorological
 905 Organization, Geneva, Switzerland, 2015.

906 World Meteorological Organization (WMO): Manual on the Global Observing
 907 System, WMO-No. 544, World Meteorological Organization, Geneva, Switzerland,
 908 2017.

909 World Meteorological Organization (WMO): Guide to Instruments and Methods of
 910 Observation, WMO-No. 8, World Meteorological Organization, Geneva, Switzerland,
 911 2018.

912 Xie, P., Chen, M., Yang, S., Yatagai, A., Hayasaka, T., Fukushima, Y., and Liu, C.: A
 913 gauge-based analysis of daily precipitation over East Asia, *J. Hydrometeorol.*, 8,
 914 607–626, <https://doi.org/10.1175/JHM583.1>, 2007.

915 Yang, S., Jones, P. D., Jiang, H., and Zhou, Z.: Development of a near-real-time global
 916 in situ daily precipitation dataset for 0000–0000 UTC, *Int. J. Climatol.*, 40,
 917 2795–2810, <https://doi.org/10.1002/joc.6367>, 2020.

918 Yang, S., Zhai, P., Zhang, X., and Zhou, Z.: Development of a global in-situ daily
 919 temperature dataset preferentially for 0000–2400 UTC from 1981–2024, Zenodo [data
 920 set], <https://doi.org/10.5281/zenodo.17895292>, 2025.

921 Yatagai, A., Kamiguchi, K., Arakawa, O., Hamada, A., Yasutomi, N., and Kitoh, A.:
 922 APHRODITE: Constructing a long-term daily gridded precipitation dataset for Asia
 923 based on a dense network of rain gauges, *Bull. Am. Meteorol. Soc.*, 93, 1401–1415,

924 <https://doi.org/10.1175/BAMS-D-11-00122.1>, 2012.

925 Zhang, X., Alexander, L., Hegerl, G. C., Jones, P., Tank, A. K., Peterson, T. C., Trewin,
926 B., and Zwiers, F. W.: Indices for monitoring changes in extremes based on daily
927 temperature and precipitation data, *Wiley Interdiscip. Rev.: Clim. Change*, 2, 851–870,
928 <https://doi.org/10.1002/wcc.147>, 2011.