

1 A 25 km Daily Gridded Dataset of Meteorological Variables and High-Impact  
2 Weather Events for New-type Power Systems in China

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23 **Abstract:** The new-type power system exhibits pronounced “weather dependency”,  
24 wherein high-impact weather events can significantly exacerbate operational security  
25 risks. A high-quality gridded dataset that involves both meteorological variables and  
26 high-impact weather events is of great significance for new-type power systems. In this  
27 study, a spatially adaptive optimal interpolation scheme is developed and applied to  
28 generate the China New-type Power Systems Meteorological (CNPS-Met) dataset. The  
29 CNPS-Met dataset covers the entire Chinese mainland, with a daily temporal resolution  
30 and a 25 km spatial resolution. It includes eight meteorological variables and eleven  
31 high-impact weather events, categorized from generation-side, grid-side and demand-  
32 side perspectives relevant to new-type power systems. Validation with existing datasets  
33 indicates that the CNPS-Met dataset generally exhibits superior performance in  
34 meteorological estimation. Specifically, the estimated mean relative errors for 2-m air  
35 temperature, 2-m specific humidity, 10-m wind speed, precipitation and surface  
36 pressure averaged over the Chinese mainland could be reduced by 1.7%-18.5%, 9.0%-  
37 29.6%, 1.9%-8.5%, 2.7%-18% and 4.9%-5.2%, respectively. On this basis, a series of  
38 high-impact weather events critical to the operation of new-type power system are  
39 identified. The spatial distribution of their frequency hotspots and intensity extremes  
40 are further analyzed. The CNPS-Met dataset is expected to benefit research and  
41 applications at the intersection of meteorology and new-type power systems.

42

43 **1. Introduction**

44 A high-quality meteorological reanalysis dataset is of great significance for  
45 analyzing climate change, verifying climate simulations, identifying high-impact  
46 weather events, and predicting future climate change etc. (Qin et al. 2022; Wen et al.  
47 2023). Over the past decades, China has built a large-scale ground-based  
48 meteorological observation network, with the total number of ground-based  
49 observation stations exceeding 2400 (Xu et al. 2019). However, in regions with  
50 complex terrain such as mountainous areas, the Tibetan Plateau, and the Gobi Desert,  
51 ground-based observation stations are relatively sparse. As a result, the climate  
52 variability at small geographic scales cannot be adequately represented (Wen et al. 2023;  
53 Jiang et al. 2023), which constrains the practical applications of ground-based  
54 observation stations. Recently, China has been building a new-type power system, with  
55 the core objective being to maximize the integration of renewable energy such as wind  
56 and solar energy (Xin 2023). However, renewable energy integration is highly  
57 susceptible to weather and climate (D'Amico et al. 2024; Gao et al. 2025). Against the  
58 backdrop of global warming and the increasing frequency of extreme weather events  
59 (IPCC, 2021), significant challenges are expected for the development of the new-type  
60 power system. Therefore, to support both research and practical needs related to new-  
61 type power systems, it is essential and urgent to develop a high-quality gridded dataset  
62 that includes both meteorological variables and high-impact weather events relevant to  
63 power systems.

64 Apart from several global atmospheric reanalysis datasets such as the ECMWF

65 (European Centre for Medium-Range Weather Forecasts) Reanalysis v5 (ERA5)  
66 (Hersbach et al. 2020), and Modern-Era Retrospective analysis for Research and  
67 Applications (MERRA) (Christensen et al. 2019) etc., several other widely used  
68 gridded meteorological datasets covering China have recently been developed, most of  
69 which are available at a daily scale. For instance, the gridded daily observation dataset  
70 across the China region (CN05.1) was developed based on approximately 2400 ground-  
71 based observation stations in China. It has a spatial resolution of  $0.25^{\circ} \times 0.25^{\circ}$  and  
72 covers the period from 1961 to 2020. This dataset was constructed using spatial  
73 interpolation methods (Wu and Gao 2013; Wu et al. 2017). The China Meteorological  
74 Forcing Data (CMFD) dataset, spanning from 1951 to 2020 with a temporal resolution  
75 of 3 h and a spatial resolution of  $0.1^{\circ} \times 0.1^{\circ}$ , was produced by integrating remote sensing  
76 products, ERA5 reanalysis, and approximately 400 ground-based observation stations  
77 in China. The methodology employed interpolation techniques based on the  
78 ANUSPLIN software and deep learning (He et al. 2020). More recently, the China Daily  
79 Meteorological Dataset (CDMet), covering 2000 to 2020, at a spatial resolution of 4  
80 km  $\times$  4 km, was generated by merging ERA5 reanalysis with 699 ground-based  
81 meteorological stations across China. An adaptive interpolation scheme combining  
82 thin-plate spline interpolation and random forest algorithm was used in its production  
83 (Zhang et al. 2024). These datasets provide useful basis for climate analysis, land  
84 surface and hydrology process study etc. (e.g., Qiu et al. 2024; Sutanto et al. 2024).  
85 Extreme weather and climate events can also be derived from these datasets, using  
86 indices released by the World Meteorological Organization (Heim et al. 2015). However,

87 the definition of extreme weather and climate events in atmospheric sciences, typically  
88 conceptualized as low-probability events under large-sample assumptions, may not  
89 fully align with the operational needs of new-type power systems. In fact, there are  
90 currently no dedicated datasets of extreme or high-impact weather events categorized  
91 according to the generation-side, grid-side, and demand-side needs of new-type power  
92 systems. Furthermore, although both the CDMet and CMFD datasets incorporate  
93 diverse data sources, including satellite remote sensing and reanalysis products, their  
94 utilization of ground-based observation stations remains relatively limited. Over the  
95 complex terrain, ground-based observation stations have been shown to possess  
96 superior accuracy and representativeness compared to satellite-derived and reanalysis  
97 data (Wei et al. 2023; Rao et al. 2024; Jiang et al. 2025).

98 Another issue that requires attention is that the methodology employed in the  
99 aforementioned datasets relies heavily on spatial interpolation techniques. When  
100 limited ground-based observation stations are used to generate gridded dataset at finer  
101 resolution, the process effectively becomes extrapolation, meaning that estimates are  
102 made beyond the boundaries of the original data coverage. In contrast, data assimilation,  
103 a well-established technique in atmospheric modelling, aims to optimally combine  
104 observations with background model fields to produce a more accurate estimate of the  
105 true atmospheric state, while explicitly accounting for uncertainties in both the  
106 observations and the model (Talagrand 1997). Additionally, data assimilation  
107 incorporates information about the influence of climate condition on the spatial  
108 distribution and relationships among meteorological variables (Kalnay 2003). In

109 practice, it has been widely used in operational numerical weather prediction and the  
110 construction of gridded datasets (e.g., Kalnay 2003; Hunt et al. 2007; Bannister 2008;  
111 Lee et al. 2013; Carrassi et al. 2018; Lindskog et al. 2019; Zhao et al. 2024). The optimal  
112 interpolation (OI) is a classical data assimilation scheme known for its high  
113 computational efficiency and reliable accuracy. It has been shown to be fundamentally  
114 equivalent to more advanced methods such as the three-dimensional variational  
115 assimilation (Gandin 1959; Akmaev 1999; Eyre et al. 2022). A key factor influencing  
116 the performance of OI is the design of the observation operator (e.g., Daley 1993;  
117 Uboldi et al. 2008; Giroto et al. 2020).

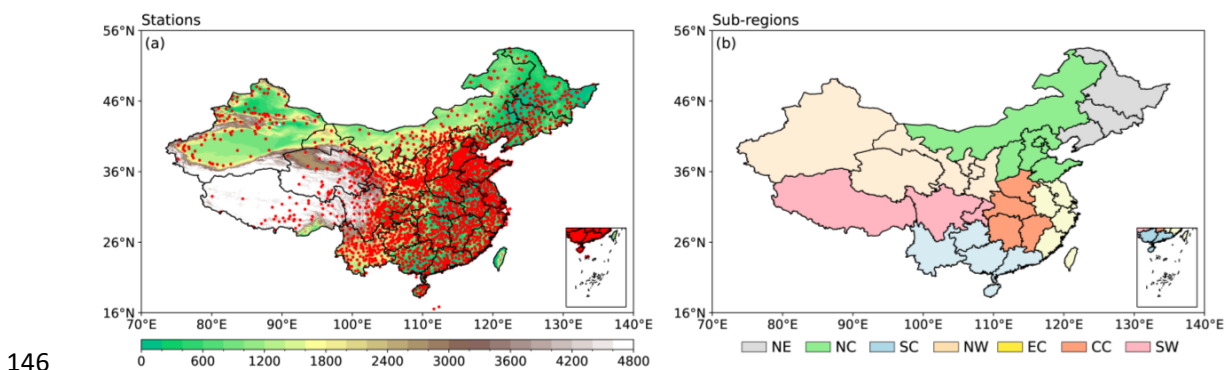
118 The Cressman interpolation method (Cressman 1959), which establishes the  
119 relationship between observations and background field through a weight function, is  
120 commonly used as observation operator in OI (Liu et al. 2016). However, in the  
121 traditional Cressman interpolation, the influence radius in the weight function is  
122 assumed to be a fixed constant. This assumption is reasonable in idealized situations  
123 where observations are uniformly distributed. In cases of uneven observational  
124 coverage, however, the use of a fixed radius can introduce significant errors and  
125 uncertainties into the observation operator, thereby degrading the performance of the  
126 OI scheme (e.g., Alonso et al. 2018; Miatselskaya et al. 2022; Wang et al. 2023; Jiang  
127 et al. 2025). Therefore, dynamically adjusting the influence radius based on the spatial  
128 distribution and density of observations around each grid point in the background field  
129 would be a potential approach to improving observation operator and enhancing the  
130 overall performance of OI. Based on the aforementioned discussions, the motivation of

131 this study is to develop an improved OI assimilation scheme, and to generate the China  
132 New-type Power Systems Meteorological (CNPS-Met) dataset. This dataset includes  
133 basic meteorological variables and high-impact weather events, categorized according  
134 to three critical vulnerability dimensions: generation-side, grid-side, and demand-side.

## 135 **2. Data and methods**

### 136 *a. Modelling data*

137 The CNPS-Met dataset is generated by fusing hourly ground-based observation  
138 stations with ERA5 reanalysis. The data from 2598 meteorological stations across  
139 China (Figure 1a), spanning the period from 1980 to 2016, are used. These data include  
140 wind speed at 10 m, air temperature, relative humidity at 2 m, surface pressure, and  
141 precipitation, and can be obtained from China Meteorological Administration  
142 (<https://data.cma.cn/>). Prior to publication, the observations underwent strict quality  
143 control. The meteorological stations are densely distributed in eastern and southern  
144 China (Fig. 1a) but are sparse in the northwestern regions and the Tibetan Plateau (Fig.  
145 1a).



146  
147 Figure 1. Distribution of (a) ground-based meteorological stations (red dots) and terrain height  
148 (shaded colors), and (b) the seven sub-regions across Chinese mainland. The seven sub-regions

149 include Northeast China (NE), North China (NC), South China (SC), Northwest China (NW), East  
150 China (EC), Central China (CC), and Southwest China (SW).

151 ERA5, the fifth generation of reanalysis data released by the ECMWF  
152 (<https://cds.climate.copernicus.eu/datasets/reanalysis-era5-land?tab=overview>),  
153 exhibits robust performance in China (Hersbach et al. 2020; Jiang et al. 2021; Lavers  
154 et al. 2022). In this study, precipitation, surface pressure, wind speed at 10 m, air  
155 temperature and specific humidity at 2 m, at a horizontal resolution of  $1^\circ \times 1^\circ$  and a  
156 temporal resolution of 1 hour, are used as background field in the assimilation. Specific  
157 humidity and relative humidity can be mutually converted through thermodynamic  
158 formulas that incorporate air temperature and pressure (Lovell-Smith et al. 2005).

159 To improve the accuracy of the input data and ensure the integrity of the CNPS-  
160 Met dataset, we exclude the anomalous records by detecting records that are deviated  
161 significantly from their mean values using the three-sigma rule method (Oakland and  
162 Oakland 2007). The three-sigma rule method was applied to the full time series.  
163 Approximately 0.18% records were excluded.

#### 164 *b. Validation data*

165 The daily CN05.1, CMFD and CDMet gridded datasets are used to validate the  
166 CNPS-Met dataset. Although the CMFD has the sub-daily (3-hourly) records, it is  
167 primarily derived from the ERA5 reanalysis and remote sensing products, rather than  
168 ground-based observation stations. Therefore, daily datasets are validated in this study.  
169 In addition, although the CMFD and CDMet have horizontal resolutions of 10 km and  
170 4 km, respectively, they are generated essentially by spatial interpolation rather than

171 fusing additional observations. Hence, all datasets are interpolated to a common  
 172 horizontal resolution of  $0.25^\circ \times 0.25^\circ$ .

173 *c. Spatially adaptive optimal interpolation assimilation scheme*

174 The Optimal Interpolation (OI) assimilation scheme is employed to generate the  
 175 CNPS-Met dataset. This scheme estimates optimal values by minimizing the errors  
 176 between the observations and the background fields. The objective function is defined  
 177 as follows:

$$178 \quad \mathbf{x}_a = \mathbf{x}_b + \mathbf{W}[\mathbf{y}_o - \mathbf{H}(\mathbf{x}_b)] \quad (1)$$

179 where  $\mathbf{x}_a$  is the analysis field (optimal field),  $\mathbf{x}_b$  is the background field (e.g.,  
 180 ERA5 reanalysis), they are both the matrix of  $m \times n$  (grid points in the latitudinal and  
 181 meridional directions, respectively);  $\mathbf{y}_o$  is the observations, which is the vector with a  
 182 length of  $p$  (e.g., number of ground-based stations); the two-dimensional matrix  $\mathbf{H}$  is  
 183 the observation operator, which maps values from regularly gridded background fields  
 184 to irregularly distributed ground-based station observations;  $\mathbf{W}$  is the optimal weight  
 185 matrix, which can be written as:

$$186 \quad \mathbf{W} = \mathbf{B}\mathbf{H}^T (\mathbf{H}\mathbf{B}\mathbf{H}^T + \mathbf{R})^{-1} \quad (2)$$

187 where superscript  $T$  denotes the matrix transpose operation;  $\mathbf{B}$  is the background  
 188 error covariance matrix, and  $\mathbf{R}$  is the observation error covariance matrix, they can be  
 189 written as:

$$190 \quad \mathbf{B} = \mathbf{E}\{\boldsymbol{\varepsilon}_b \boldsymbol{\varepsilon}_b^T\} \quad (3)$$

$$191 \quad \mathbf{R} = \mathbf{E}\{\boldsymbol{\varepsilon}_o \boldsymbol{\varepsilon}_o^T\} \quad (4)$$

192 where  $\boldsymbol{\varepsilon}_b$  is the vector of grid points variances and covariances in the background

193 filed over a given period (e.g., one month), while  $\boldsymbol{\varepsilon}_0$  is the corresponding vector of  
 194 variances and covariances for ground-based station observations over the same period;  
 195  $\mathbf{E}$  represents a two-dimensional matrix. From the above equations, it is clear that given  
 196 the observations ( $\mathbf{y}_0$ ) and the background field ( $\mathbf{x}_b$ ), the background error covariance  
 197 matrix ( $\mathbf{B}$ ) and the observation error covariance matrix ( $\mathbf{R}$ ) are determined.  
 198 Consequently, the performance of the OI assimilation scheme depends solely on the  
 199 observation operator ( $\mathbf{H}$ ).

200 The observation operator ( $\mathbf{H}$ ), implemented here using Cressman interpolation,  
 201 applies a distance-dependent weighting function to compute a weighted average of  
 202 observations, with weights monotonically decreasing as a function of distance, thereby  
 203 emphasizing the contribution of local observations to the final interpolated field. The  
 204 observation operator can be determined via iterative updating as follows:

$$205 \quad \mathbf{H}^\gamma = \frac{\sum_{k=1}^K (w_{ijk}^2 \Delta \alpha_k^\gamma)}{\sum_{k=1}^K w_{ijk}} \quad (5)$$

206 where  $\Delta \alpha_k^\gamma = \mathbf{y}_0(k) - \mathbf{x}_b^\gamma$  denotes the difference between observation at  $k^{\text{th}}$   
 207 ground-based station and grid point ( $i, j$ ) at  $\gamma^{\text{th}}$  iteration;  $K$  denotes the number of total  
 208 ground-based stations ;  $\mathbf{x}_b^\gamma = \mathbf{x}_b^{\gamma-1} + \mathbf{H}^{\gamma-1}$  denotes updated temporary background  
 209 filed at  $\gamma^{\text{th}}$  iteration, which will be used to continuously update  $\Delta \alpha_k^\gamma$  and  $\mathbf{H}^\gamma$ , the  
 210 ERA5 reanalysis will be used as first guess in the iteration; the iteration termination  
 211 condition is  $|\Delta \alpha_k^\gamma| \leq 1 \times 10^{-6}$ , the resulting  $\mathbf{H}^\gamma$  will be then used as the definitive  
 212 observation operator ( $\mathbf{H}$ ) in Eqs. (1-2) to perform OI assimilation;  $w_{ijk}$  is the weight  
 213 function in Cressman interpolation, its expression can be written as:

$$w_{ijk} = \begin{cases} \frac{R_c^2(i, j) - d_{ijk}^2}{R_c^2(i, j) + d_{ijk}^2}, & d_{ijk} \leq R_c(i, j) \\ 0 & , \quad d_{ijk} > R_c(i, j) \end{cases} \quad (6)$$

where  $d_{ijk}$  represents the spatial distance between grid point  $(i, j)$  and observation at  $k^{\text{th}}$  ground-based station;  $R_c(i, j)$  represents the influence radius.

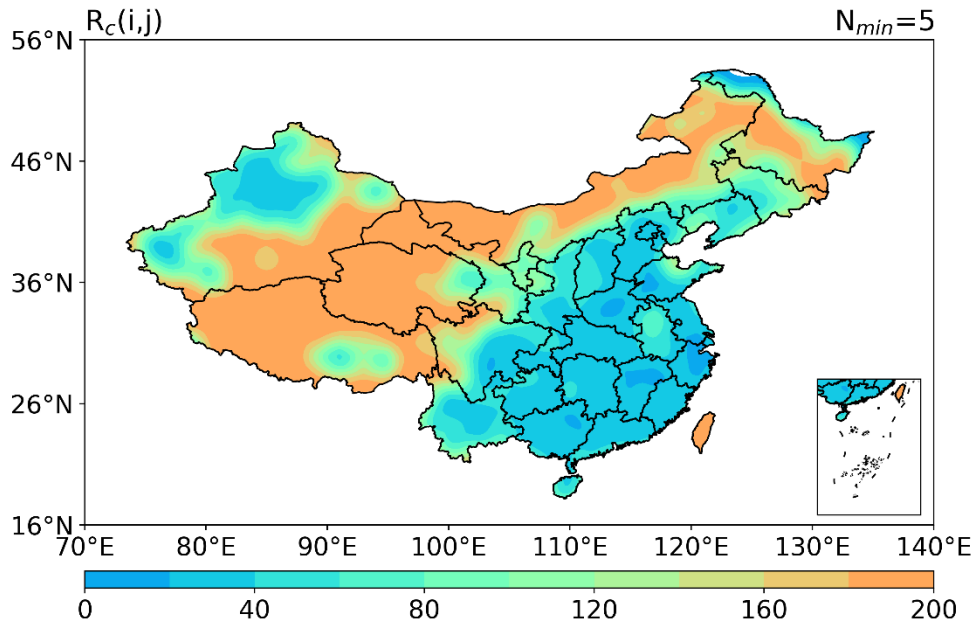
In the traditional Cressman interpolation, the influence radius is typically held constant. While this assumption is reasonable in regions with uniformly distributed observation stations, it would become problematic in practice due to the inherently uneven distribution of stations, especially over complex terrain. Such non-uniformity can degrade the performance of Cressman interpolation (Lin et al. 2012; Wang et al. 2023), and consequently impair the accuracy of OI assimilation scheme. To overcome this limitation, this study introduces a spatially adaptive influence radius that adjusts according to local observation density and distribution. This enhancement would improve the observation operator and optimizes the overall OI assimilation framework. The proposed method is referred to as the spatially adaptive OI assimilation scheme. The spatially varying influence radius  $R_c(i, j)$  is calculated as follows:

$$R_c(i, j) = \min \left\{ R \mid \hat{K}(i, j, R) \geq N_{\min}, R_{\min} \leq R \leq R_{\max} \right\} \quad (7)$$

where  $\hat{K}(i, j, R)$  denotes the number of observation stations within a circle of search radius  $R$  centered at grid point  $(i, j)$ ; the lower limit  $R_{\min}$  is 1 km, while the upper limit  $R_{\max}$  is set to 200 km;  $N_{\min}$  represents the preset minimum threshold for the number of observation stations within the search radius  $R$ . Here, this parameter is set to  $N_{\min} = 5$ , meaning that for each grid point, the scheme dynamically expands the search radius until the number of available observation stations within the search region

235 reaches at least 5. From Eq. (7), it is clear that when  $N_{min} < 5$ , in extremely data-sparse  
236 regions (e.g., Northwest China), the search radius remains too small, which may cause  
237 assimilation results based on only a few stations (e.g., 1-2 stations) to become not robust  
238 due to insufficient representativeness or accidental errors. When  $N_{min} > 5$ , this could  
239 lead to missing values of the influence radius in the data-sparse regions (not shown).

240 Figure 2 shows the spatial distribution of the influence radius  $R_c(i, j)$  in the  
241 spatially adaptive OI assimilation scheme across China. Results indicate that, the  
242 influence radius varies with the station density, that is, it is larger in data-sparse regions  
243 and is smaller in data-dense regions, which generally captures the spatial distribution  
244 of stations (Fig. 1a), suggesting that the spatially adaptive OI scheme proposed in this  
245 study could dynamically adjust the influence radius based on the density of local  
246 observations.



247

248 Figure 2. Spatial distribution of the influence radius  $R_c(i, j)$  (unit: km) in the spatially adaptive OI

249 assimilation scheme.

250 The assimilation performance of the new scheme and the traditional scheme is  
 251 compared over the sample period from January to December 2013 (not shown). Results  
 252 show that, compared with the traditional OI scheme (using a fixed influence radius),  
 253 the new scheme proposed in this study (using a spatially adaptive influence radius)  
 254 could obviously reduce the simulation errors for different regions, different months, and  
 255 different meteorological variables across China. This indicates that the new scheme  
 256 proposed in this study outperforms the traditional scheme.

#### 257 *d. Evaluation Metrics*

258 The performance of the CNPS-Met dataset is evaluated using the statistics  
 259 including the mean relative error (*MRE*), the root mean square error (*RMSE*),  
 260 correlation coefficient ( $R^2$ ), and the modeling efficiency (*EF*):

$$261 \quad MRE = \frac{1}{n} \sum_{i=1}^n |(P_i - O_i) / O_i| \quad (8)$$

$$262 \quad RMSE = \sqrt{\frac{1}{n} \sum_{i=1}^n (P_i - O_i)^2} \quad (9)$$

$$263 \quad R^2 = \frac{\left[ \sum_{i=1}^n (O_i - \bar{O})(P_i - \bar{P}) \right]^2}{\sum_{i=1}^n (O_i - \bar{O})^2 \sum_{i=1}^n (P_i - \bar{P})^2} \quad (10)$$

$$264 \quad EF = 1 - \frac{\sum_{i=1}^n (P_i - O_i)^2}{\sum_{i=1}^n (O_i - \bar{O})^2} \quad (11)$$

265 where  $n$  denotes sample size;  $O_i$  and  $P_i$  are the observed and estimated values,  
 266 respectively;  $\bar{O}$  and  $\bar{P}$  are the average of the observed and estimated values,  
 267 respectively. Values of *MRE* and *RMSE* closer to 0, and  $R^2$  and *EF* closer to 1,

268 indicate better estimation performance

269 Apart from the above statistics, a more comprehensive statistic referred to as the  
270 global performance index (*GPI*; Despotovic et al. 2015), is introduced in this study:

$$271 \quad GPI = \sum_{k=1}^4 \alpha_k (\bar{y}_k - y_k) \quad (12)$$

272 where  $\bar{y}_k$  represents the median of the scaled values of indicator  $k$  (i.e., *MRE*,  
273 *RMSE*,  $R^2$  and *EF*);  $y_k$  is scaled value of indicator  $k$ ;  $\alpha_k = 1$  corresponds to  
274 *MRE* and *RMSE*, while  $\alpha_k = -1$  corresponds to  $R^2$  and *EF*. The higher the *GPI*,  
275 the better performance of the overall estimation.

#### 276 *e. Identification of high-impact weather events for new-type power systems*

277 Based on a comprehensive review of the existing literatures, the high-impact  
278 weather events for the generation-side, grid-side and demand-side of new-type power  
279 systems could be defined in Table 1. In the generation-side, cut-out wind speed is  
280 defined as hourly wind speed reaches or exceeds  $25 \text{ m s}^{-1}$ , that is, wind turbine  
281 automatically shuts down to prevent equipment damage when wind speeds reach or  
282 exceed this threshold, resulting in an abrupt reduction of wind power output to zero  
283 (Jerez et al. 2015; Song et al. 2022). According to Jerez et al. (2015) and Song et al.  
284 (2022), cut-in wind speed is defined as hourly mean wind speeds  $\leq 2.5 \text{ m s}^{-1}$ , that is,  
285 wind turbines would remain in standby or idle mode when wind speed is less than or  
286 equal to this threshold, resulting in effectively zero power output. The wind turbine hub  
287 height defined in this study is 70 m, when identifying cut-in and cut-out wind speed  
288 that are relevant to high-impact weather events, the wind speeds at 10 m are converted  
289 to 70 m using the empirical power law method, which can be expressed as:

290 
$$u_2 = u_1 \left( \frac{h_2}{h_1} \right)^\alpha \quad (13)$$

291 where  $u_2$  and  $u_1$  represent wind speed at 70 m and 10 m, respectively;  $h_2$  and  
 292  $h_1$  represent the target height (70 m) and the reference height (10 m), respectively;  $\alpha$   
 293 is a prescribed constant, taken as 0.14.

294 Based on the observations of hourly solar irradiance and power generation  
 295 efficiency in large-scale photovoltaic power plants, Sundaram et al. (2024)  
 296 demonstrated that photovoltaic conversion efficiency decreases significantly when  
 297 hourly solar irradiance falls below  $100 \text{ W m}^{-2}$ , with the performance ratio declining to  
 298 critical levels; supporting this finding, Lei et al. (2025) established through  
 299 comprehensive literature reviews that  $\leq 100 \text{ W m}^{-2}$  represents the standardized  
 300 threshold for low-light conditions in photovoltaic systems; therefore, low radiation is  
 301 defined as hourly solar irradiance  $\leq 100 \text{ W m}^{-2}$ . Through systematic analysis of  
 302 measurements and experiments (Oloufemi et al. 2016; Mohammad et al. 2021; Yang et  
 303 al. 2022; Sun et al. 2022; Ju et al. 2022; Köster et al. 2023), Bi et al. (2025) derived a  
 304 fitted relationship between power generation loss and air temperature; for operational  
 305 definitions, extreme high temperature is specified as  $\geq 35 \text{ }^\circ\text{C}$ , while extreme low  
 306 temperature is defined as  $\leq -20 \text{ }^\circ\text{C}$ .

307 Table 1. Classification and definition of high-impact weather events for new-type power systems.

Components of new-type power system	High-impact weather events	Abbreviation	Definition	Impacts on new-type power systems	References
Generation-side	Cut-out wind speed	V <sub>out</sub>	Hourly wind speed $\geq 25 \text{ m s}^{-1}$	Wind turbine shutdown causes abrupt drop in wind power output to	Song et al. (2022) Jerez et al. (2015)

	Cut-in wind speed	V <sub>in</sub>	Hourly wind speed $\leq 2.5 \text{ m s}^{-1}$	zero Wind turbine remains in standby or idle mode, resulting in abnormal zero power output	Song et al. (2022) Jerez et al. (2015)
	Low radiation	Lowrad	Hourly radiation $\leq 100 \text{ W m}^{-2}$	Reduces the efficiency of photovoltaic conversion	Sundaram et al. (2024) Lei et al. (2024)
	Extreme high temperature	T <sub>maxg</sub>	Hourly temperature $\geq 35 \text{ }^\circ\text{C}$	Overloading of power equipment leads to loss of power generation efficiency	Mohammad et al. (2021) Yang et al. (2022)
	Extreme low temperature	T <sub>ming</sub>	Hourly temperature $\leq -20 \text{ }^\circ\text{C}$	Equipment shutdown resulting in loss of power generation efficiency	Ju et al. (2022) Sun et al. (2022)
	Ice accretion	Icing	Hourly temperature $\leq 0 \text{ }^\circ\text{C}$ , hourly relative humidity $\geq 85\%$ , and hourly wind speed $\leq 4 \text{ m s}^{-1}$ simultaneously	Significantly increases the mechanical load on transmission lines, causing line breakage, flashover, and tripping	Gu et al. (2010) Shen et al. (2010) Pei et al. (2024)
Grid-side	Snowfall	Snowing	Hourly precipitation $\geq 0.1 \text{ mm}$ and hourly temperature $\leq 0 \text{ }^\circ\text{C}$ simultaneously	Increases the risk of line icing, damages the structural strength of power facilities, and threatens the reliability of power supply	Iver et al. (2019) Wesley et al. (2020)
	Conductor galloping	Galloping	Hourly relative humidity $\geq 75\%$ and wind speeds exceeding $4 \text{ m s}^{-1}$ persisted for more than 3 hours simultaneously	Cause short circuit tripping of the line and may lead to chain faults	Tsujimoto et al. (1983) Li et al. (2015)
Demand-side	Extreme high temperature	T <sub>maxd</sub>	Hourly temperature $\geq 38 \text{ }^\circ\text{C}$	The demand for electricity load would sharply increase	Fu et al. (2015) Ye et al. (2024)
	Extreme low temperature	T <sub>mind</sub>	Hourly temperature $\leq -10$	The sensitivity of electricity	Blake et al. (2022)

		°C	load demand would sharply increase to extreme low temperature	Millin et al. (2024)
Heat and humid environment (High enthalpy environment)	HHE	Hourly temperature $\geq 28$ °C and relative humidity $\geq 65\%$ simultaneously	Significantly increases the risk of human heat stress and exacerbates the load on power equipment	Patrick et al. (2015) Jane et al. (2023)

308 In the grid-side, ice accretion is defined as hourly air temperature  $\leq 0$  °C, hourly  
309 relative humidity  $\geq 85\%$  and hourly wind speed  $\leq 4$  m s<sup>-1</sup>; this definition is supported  
310 by three evidences: first, thermodynamic analysis by Gu et al. (2010) demonstrated  
311 through thermal equilibrium theory and wind tunnel experiments that the required Joule  
312 heating for anti-icing systems exhibits a sharp decline when temperatures fall below  
313 0 °C, indicating a fundamental threshold for ice formation; second, comprehensive field  
314 observations by Shen et al. (2010) established the multi-parameter requirements for ice  
315 accretion on transmission lines, that are, the critical thermal window (temperature  $\leq 0$   
316 °C, with optimal range between -10 °C and -1 °C), the moisture threshold (relative  
317 humidity  $\geq 85\%$  for sufficient water vapor supply), and the aerodynamic constraint  
318 (wind speed  $\leq 4$  m s<sup>-1</sup> to enable effective droplet impingement while preventing wind-  
319 driven shedding); third, these parameters are also codified in the Chinese  
320 Meteorological Industry Standard QX/T 355-2016 for wire icing risk assessment, which  
321 formally defines ice accretion as “the adherence of glaze, rime, or frozen wet snow to  
322 conductors” (Pei et al. 2024). Tsujimoto et al. (1983) found that conductor galloping  
323 typically occurs when wind speeds  $\geq 4$  m s<sup>-1</sup> and persist for over 3 hours; Li et al. (2015)  
324 further established meteorological thresholds by analyzing hourly weather variations

325 during galloping events and considering galloping mechanisms and grid operation  
326 experience; based on these studies, the galloping criterion in this study is defined as:  
327 hourly relative humidity  $\geq 75\%$  with sustained ( $\geq 3$  h) wind speeds  $\geq 4 \text{ m s}^{-1}$ . Snowfall  
328 is defined as hourly precipitation  $\geq 0.1 \text{ mm}$  with air temperature  $\leq 0 \text{ }^\circ\text{C}$ , consistent  
329 with the standard definition adopted in community land surface models (Oleson et al.  
330 2013).

331 In the demand-side, Fu et al. (2015) investigated the response of observed daily  
332 peak power load to temperature variations, identifying  $38 \text{ }^\circ\text{C}$  as a critical threshold for  
333 peak power load, beyond which demand surges dramatically; observation analysis of  
334 Shaffer et al. (2022) found that power demand sensitivity increases sharply below  $-10$   
335  $^\circ\text{C}$ ; similarly, Millin et al. (2024) observed significant load anomalies below  $-6 \text{ }^\circ\text{C}$  in  
336 the U.S. Midwest; accordingly, we define extreme high and low temperature thresholds  
337 as: hourly temperature  $\geq 38 \text{ }^\circ\text{C}$  and  $\leq -10 \text{ }^\circ\text{C}$ , respectively. Baldwin et al. (2023)  
338 demonstrated through physiological experiments and observations that combined  
339 thermal stress (air temperature  $\geq 30 \text{ }^\circ\text{C}$  with relative humidity  $\geq 65\%$ ) significantly  
340 increases human heat strain risks in power load sectors; Sullivan et al. (2015) further  
341 identified  $28 \text{ }^\circ\text{C}$  as the critical temperature threshold for notable load growth through  
342 hourly load-temperature analysis; accordingly, heat and humid environment (high  
343 enthalpy environment) is defined as: hourly temperature  $\geq 28 \text{ }^\circ\text{C}$  with relative  
344 humidity  $\geq 65\%$ .

345 We need to explain that although these high-impact weather events are defined  
346 through literature reviews, their definitions are grounded in empirical evidence derived

347 from observational studies, controlled laboratory experiments, or synthesis of  
348 established research findings. Therefore, the resulting classifications should be both  
349 scientifically reasonable and reliable. Furthermore, the CNPS-Met dataset is generated  
350 by assimilating hourly *in-situ* observations into hourly ERA5 reanalysis; therefore, the  
351 minimum temporal resolution of the meteorological variables is 1 hour. On this basis,  
352 high-impact weather events are identified according to their respective definitions.  
353 After all such events are identified at hourly scale; they are aggregated to the daily scale.  
354 In other words, the CNPS-Met dataset supports both hourly and daily temporal scales.  
355 The hourly variables, including all meteorological elements and high-impact weather  
356 events, are subsequently stored and published online at daily scale. Moreover, the  
357 “frequency” in the following text refers to the number of days where the event occurred  
358 (i.e. the number of days where the event occurred at least once), rather than an estimate  
359 of the number of hours. For example, if a grid point experiences a high-impact weather  
360 event for at least one hour on a certain day, then that day is marked as a high-impact  
361 weather event day for that grid point.

362 For high-impact weather events such as ice accretion, conductor galloping, and  
363 heat and humid environment in Table 1, as they involve multiple meteorological  
364 variables, the following composite weather index (*CWI*) is defined to characterize their  
365 occurrence and intensity:

$$366 \quad CWI = \begin{cases} \prod_{k=1}^n \frac{\alpha_k - th(\alpha_k)}{\max(\alpha_k) - th(\alpha_k)}, & \alpha_1 \geq th(\alpha_1), L, \alpha_n \geq th(\alpha_n) \\ 0 & , else \end{cases} \quad (14)$$

367 where  $\alpha$  represents a high-impact weather event composed of  $n$  meteorological

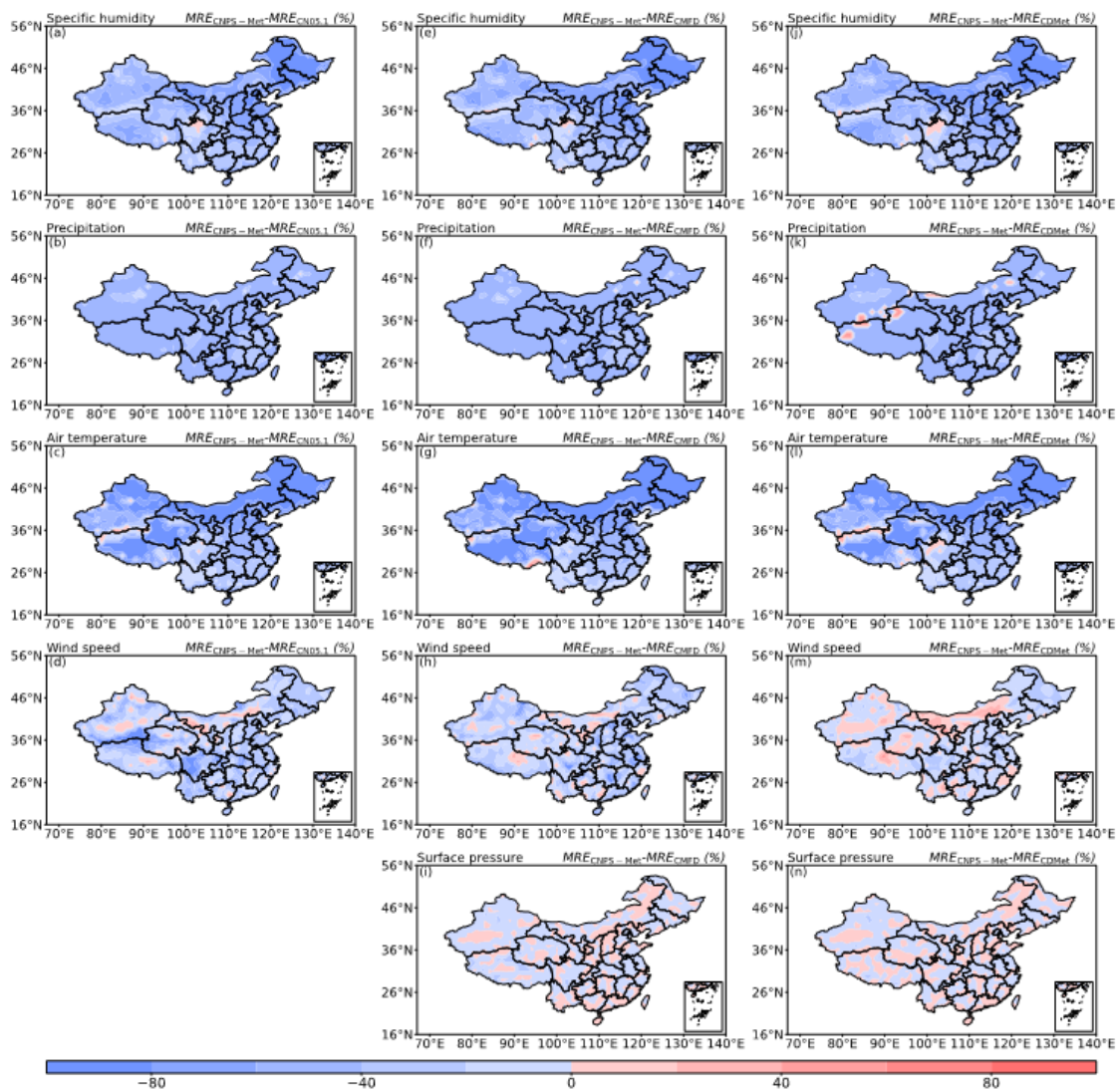
368 variables, where the index of each variable is denoted by subscript  $k$  ( $k=1, 2, \dots, n$ ).  
369 The threshold and the daily maximum value of the  $k$ -th variable ( $\alpha_k$ ) are denoted as  
370  $th(\alpha_k)$  and  $\max(\alpha_k)$ , respectively. The  $\max(\alpha_k)$  represents the multi-year daily  
371 maximum value of the  $k$ -th variable in the corresponding different grid point.

372 To analyze high-impact weather events affecting new-type power systems across  
373 different regions of China, seven sub-regions (Fig. 1b) are defined according to the  
374 spatial distribution and organizational characteristics of the power grid in China (Zhuo  
375 et al. 2022).

### 376 **3. Verification of the CNPS-Met dataset**

377 Figure 3 shows the spatial distribution of differences in *MREs* of various  
378 meteorological variables between the CNPS-Met dataset and three other widely used  
379 datasets (CN05.1, CMFD and CDMet). Results show that the CNPS-Met dataset  
380 achieves lower *MREs* across different meteorological variables and over the majority  
381 region of China compared to the other datasets, indicating a generally higher accuracy  
382 of the meteorological estimates in CNPS-Met. Significant improvements are  
383 particularly evident in humidity, temperature and precipitation. However, exceptions  
384 are observed in some regions along the periphery of the Tibetan Plateau, where  
385 performance gains are less pronounced. Compared to the other datasets, the  
386 improvement in wind speed within CNPS-Met remains limited. Consistent results can  
387 also be found in different seasons (not shown). These discrepancies may be attributed  
388 to the following factors. First, the OI assimilation scheme employed in this study relies  
389 on background and observation error covariance matrices [Eqs. (3-4)] derived from

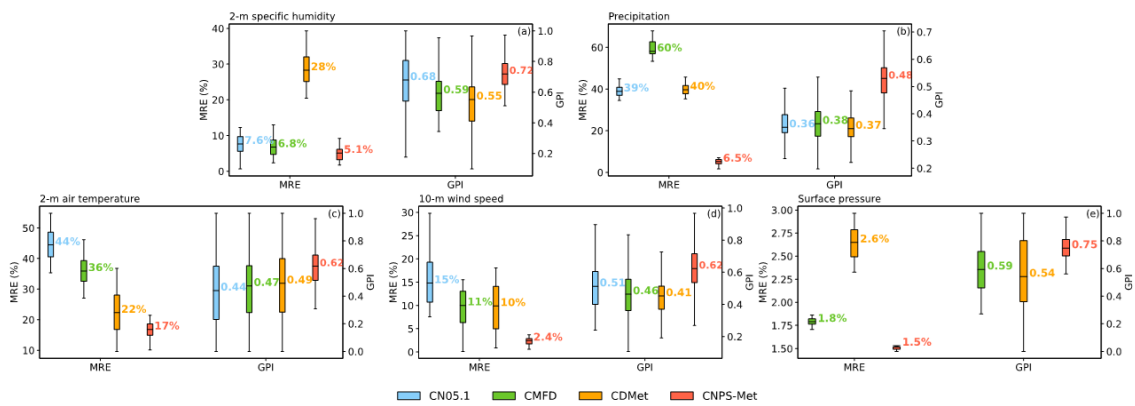
390 monthly-scale statistics. These matrices are static and may fail to adequately capture  
 391 the rapid temporal variation characteristics of highly transient and intermittent variables  
 392 such as wind speed. Second, regions where CNPS-Met exhibits larger errors are  
 393 characterized by complex terrain and sparse observational coverage, the inherent  
 394 uncertainties in the background field (e.g., ERA5) would diminish the effectiveness of  
 395 the assimilation performance in these regions.



396  
 397 Figure 3. Spatial distribution of the differences in the mean *MREs* (unit: %; averaged over 1980-  
 398 2016) between three dataset pairs: (a-d) CNPS-Met and CN05.1 ( $MRE_{\text{CNPS-Met}} \text{ minus } MRE_{\text{CN05.1}}$ ),  
 399 (e-i) between CNPS-Met and CMFD ( $MRE_{\text{CNPS-Met}} \text{ minus } MRE_{\text{CMFD}}$ ), and (j-n) between CNPS-Met

400 and CDMet ( $MRE_{CNPS-Met}$  minus  $MRE_{CDMet}$ ). The differences are shown for (a, e, j) 2-m specific  
 401 humidity, (b, f, k) precipitation, (c, g, l) 2-m air temperature, (d, h, m) 10-m wind speed, and (i, n)  
 402 surface pressure. Note that CN05.1 dataset does not include surface pressure.

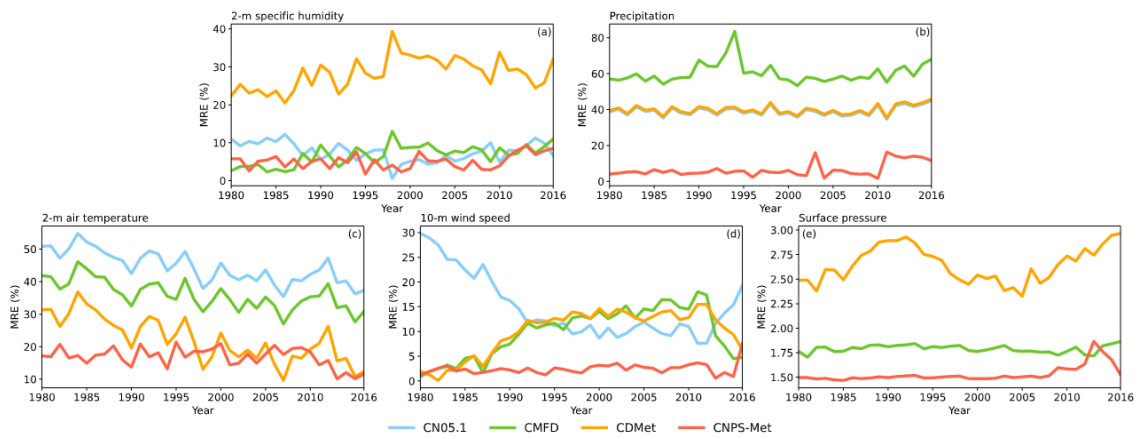
403 Figure 4 displays box plots of the  $MREs$  and  $GPI$  values across different datasets  
 404 and meteorological variables, averaged over China for the period 1980-2016. In  
 405 comparison to the other datasets, CNPS-Met exhibits the lowest  $MREs$  with the  
 406 narrowest range. Similarly, the  $GPI$  values in CNPS-Met are generally closest to 1.0  
 407 and show lower variability among the datasets. These results collectively indicate that  
 408 the CNPS-Met dataset achieves superior performance over existing alternatives.



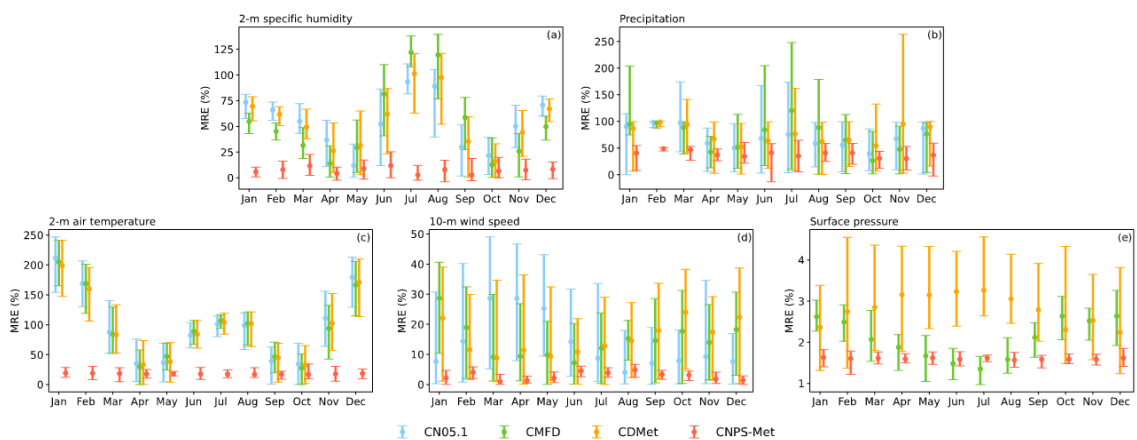
409  
 410 Figure 4. The mean  $MREs$  (unit: %) and  $GPIs$  (unit: dimensionless) averaged over China from 1980  
 411 to 2016 in different datasets for (a) 2-m specific humidity, (b) precipitation, (c) 2-m air temperature,  
 412 (d) 10-m wind speed, and (e) surface pressure.

413 To evaluate the effects of CNPS-Met at temporal scale, Figure 5 compares the  
 414 annual variations of  $MREs$  in China for different meteorological variables across  
 415 different datasets. Results show that CNPS-Met generally outperforms other datasets in  
 416 most years, especially for precipitation, wind speed and surface pressure. Exceptions  
 417 occur for air temperature and specific humidity, where  $MREs$  from CNPS-Met are larger,

418 such as near 1985 and between 2005 and 2010. The monthly *MREs* across different  
 419 datasets and meteorological variables, averaged over China for the period 1980-2016,  
 420 are further compared in Figure 6. Consistent with the above results, CNPS-Met  
 421 outperforms the other datasets in different months, exhibiting generally the lowest  
 422 *MREs* and narrowest variability range. As noted earlier, the improvement effect of  
 423 CNPS-Met on precipitation remains modest compared to that on other meteorological  
 424 variables.



425  
 426 Figure 5. The inter-annual variation of the mean *MREs* (unit: %; averaged over China) for (a) 2-m  
 427 specific humidity, (b) precipitation, (c) 2-m air temperature, (d) 10-m wind speed and (e) surface  
 428 pressure in different datasets.



429

430 Figure 6. Monthly variation of the mean *MREs* (unit: %; averaged in China from 1980 to 2016) for  
431 (a) 2-m specific humidity, (b) precipitation, (c) 2-m air temperature, (d) 10-m wind speed and (e)  
432 surface pressure in different datasets.

433 Given the apparent spatial heterogeneity of *MREs* across different datasets (Fig.  
434 3), Figure 7 presents the *MREs* averaged over the period from 1980 to 2016 for China  
435 and its seven sub-regions. Results show that among all datasets evaluated, CNPS-Met  
436 demonstrates the lowest *MREs* in various meteorological variables over both the entire  
437 China region and its seven sub-regions. In addition to the findings consistent with the  
438 analysis above, that are, the *MREs* for different meteorological variables in CNPS-Met  
439 are the smallest. Compared to the other three datasets, *MREs* of air temperature, specific  
440 humidity, wind speed, precipitation and surface pressure averaged over China for the  
441 past 40 years could be reduced by 1.7%-18.5%, 9.0%-29.6%, 1.9%-8.5%, 2.7%-18%  
442 and 4.9%-5.2%, respectively. For specific humidity, CNPS-Met exhibits relatively  
443 small *MREs* (7-9%) in South China (SC), East China (EC), Central China (CC), and  
444 Northeast China (NE), whereas relatively large *MREs* (approximately 20%) are  
445 observed in Northwest China (NW) and Southwest China (SW). For wind speed, the  
446 smallest *MRE* (4.1%) occurs in Northeast China (NE), while the largest *MRE* (9.0%) is  
447 found in North China (NC). In the case of air temperature, smaller *MREs* (below 3%)  
448 are exhibited in East China (EC) and Central China (CC), contrasting with the largest  
449 *MREs* (14.1%) in Northwest China (NW). For precipitation, the smallest *MRE* (9.6%)  
450 is observed in Northwest China (NW), compared to the largest *MRE* (57.8%) in East  
451 China (EC). For surface pressure, the smaller *MRE* (below 10%) occurs in Northeast

452 China (NE), North China (NC), Central China (CC), South China (SC) and East China  
 453 (EC), while the larger *MRE* (9.0%) is found in other regions. Noted that the  
 454 improvement of CNPS-Met in wind speed is relatively modest compared to other  
 455 datasets (see Figs. 3). However, wind speed in CNPS-Met exhibits the smallest *MREs*  
 456 among all meteorological variables, similar phenomenon can also be observed in other  
 457 datasets (see Figs. 5-7).

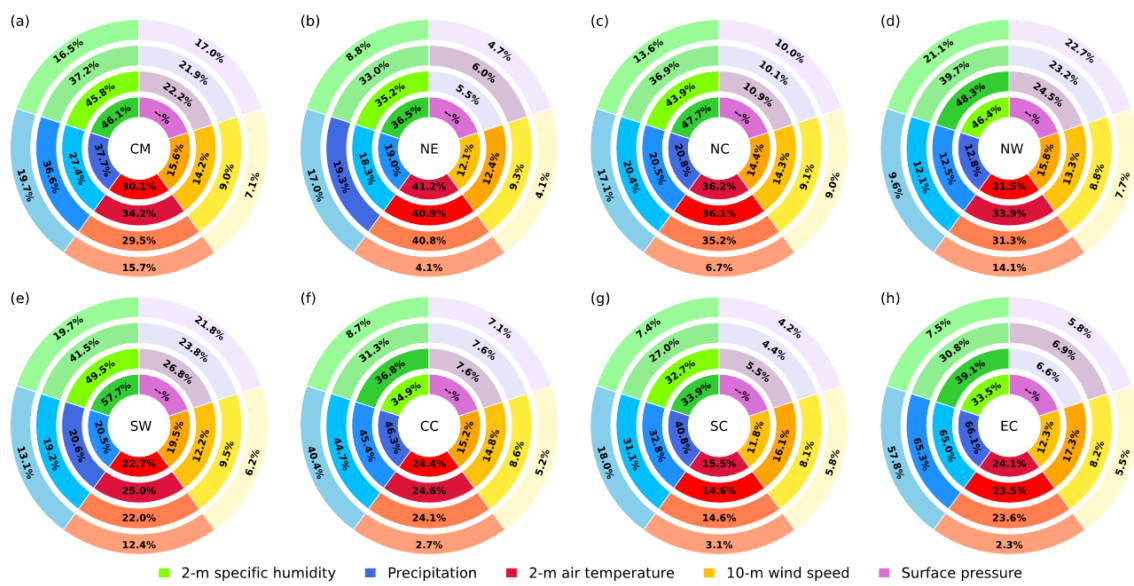


Figure 7. The mean *MREs* (unit: %; averaged over 1980-2016) of different meteorological variables in (a) Chinese mainland (CM), (b) Northeast China, (c) North China, (d) Northwest China, (e) Southwest China, (f) Central China, (g) South China, and (h) East China. The concentric circles represent different datasets (from inner to outer: CN05.1, CMFD, CDMet and CNPS-Met). The lowest values of *MREs* are denoted as the lightest color. The mean *MREs* for surface pressure are denoted as --%, as it is not included in the CN05.1 dataset.

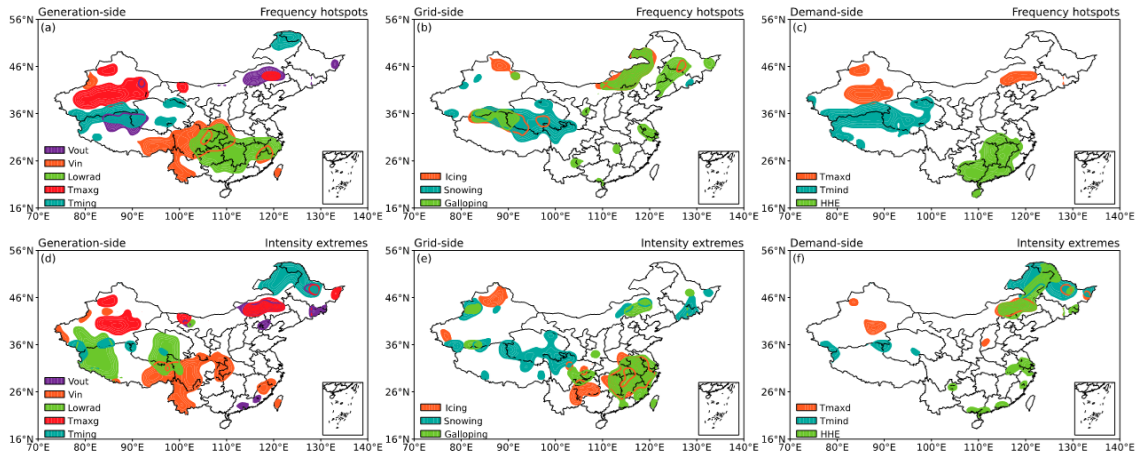
#### 4. Characteristics of high-impact weather events for new-type power systems

In this section, high-impact weather events from three critical dimensions of the new-type power system such as generation-side, grid-side, and demand-side will be

468 identified from Table 1, followed by a discussion of their spatiotemporal characteristics  
469 in the past 40 years.

470 Figure 8 shows the spatial distribution of the multi-year averaged frequency  
471 hotspots and intensity extremes (90% confidence level) of different high-impact  
472 weather events in China. The “intensity extreme” at the 90% confidence level is  
473 obtained through T-test and refers to the 90<sup>th</sup> percentile of intensity of high-impact  
474 weather events. In the generation-side, cut-out wind speed predominantly occurs over  
475 the northern Tibetan Plateau, Eastern Inner Mongolia, and parts of Xinjiang known as  
476 the “Hundred-mile Wind Zone”, which is consistent with the regions of high wind  
477 energy potentials, as analyzed by Pan et al. (2012), Yao et al. (2018) and Gyatso et al.  
478 (2023). Cut-in wind speed is primarily observed in Southwest China, this spatial pattern  
479 aligns with existing research on sustained weak wind events in Chinese Mainland,  
480 which are known to severely impact generation-side reliability (Gao et al. 2025). Low  
481 radiation events are concentrated in the middle and lower reaches of the Yangtze River.  
482 This finding is consistent with Zhang et al. (2024), who attribute the region’s lower  
483 solar radiation to its higher cloud cover and humidity. Extreme high temperatures are  
484 primarily found in the desert regions of Xinjiang (i.e., Junggar and Tarim basins), as  
485 well as in Eastern Inner Mongolia, a pattern highly consistent with existing climate  
486 model simulation and observations and largely attributed to regional arid conditions  
487 (Meng et al. 2019; Dong et al. 2024). Extreme low temperatures occur most frequently  
488 in the Kunlun Mountains, the Qilian Mountains and Northeast China, which is  
489 consistent with Yang et al. (2015) and Shi et al. (2016), who note that despite a general

490 decline trend of extreme low temperatures, these regions remain prone to such events.  
491 In the grid-side, ice accretion primarily affects Northeast China, Northern Xinjiang and  
492 Kunlun Mountains, which is also reported by Chen et al. (2010). Snowfall events are  
493 most frequent across the Tibetan Plateau, Northeast China, and Northwest Xinjiang,  
494 this distribution pattern is consistent with the findings of Yang et al. (2019) and Wang  
495 et al. (2022) based on their analysis of observations and multi-source reanalysis datasets.  
496 Conductor galloping occurs mainly in Northeast China, northern Tibetan Plateau, and  
497 sporadic regions in southern China. The spatial distributions of extreme high- and low-  
498 temperature frequencies in the demand-side are similar to those in the generation-side.  
499 Heat and humid environments occur primarily in Central and Southern China,  
500 consistent with Li et al. (2025) regarding their impact on the demand-side. The spatial  
501 distributions of high-impact weather intensity and frequency are generally consistent,  
502 albeit with some exceptions. For example, in the generation-side, low solar radiation  
503 events are most frequent in the middle and lower reaches of the Yangtze River, yet they  
504 are relatively weak when they occur. In the grid-side, ice accretion is infrequent in  
505 Southern China but tends to be intense. In the demand-side, the extreme low  
506 temperatures in Northeast China are particularly severe.



507

508 Figure 8. Spatial distribution of frequency hotspots and intensity extremes (90% confidence level)

509 of different high-impact weather events in Chinese mainland during 1980-2016.

510 Figures 9-11 summarize the frequency and intensity of high-impact weather events

511 in the generation-side, grid-side and demand-side in China and its sub-regions. In the

512 generation-side, the highest frequency of cut-out wind speed occurs in North China,

513 while its highest intensity is in East China. Cut-in wind speed is most frequent in

514 Southwest and Central China. Low radiation occurs most frequently in East and Central

515 China. Extreme high temperatures are relatively frequent in Northwest, Central, East

516 and South China, with the greatest intensity observed in North China. Extreme low

517 temperatures are most frequent and most intense in Northeast China. On average, the

518 frequency and mean intensity of cut-out wind speed, cut-in wind speed, low radiation,

519 extreme high temperature and extreme low temperature in China are 0.4% and 37.3 m

520  $s^{-1}$ , 58.9% and  $1.5 m s^{-1}$ , 14.9% and  $30.1 W m^{-2}$ , 2.5% and  $37.1 ^\circ C$ , 9.9% and  $-23.1 ^\circ C$ ,

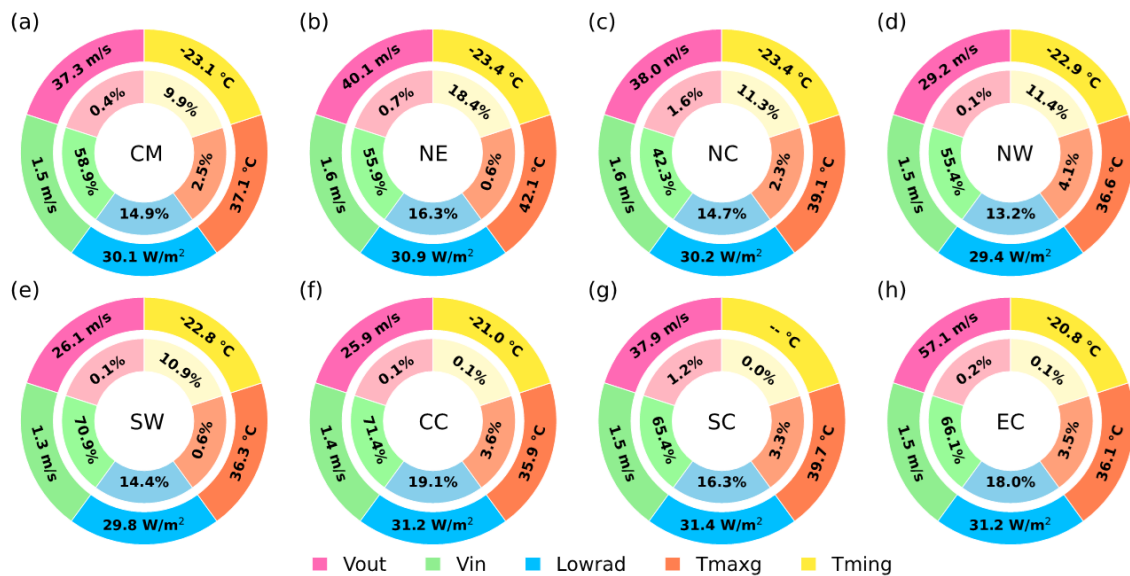
521 respectively. In the grid-side, ice accretion occurs most frequently in North China while

522 its most severe events are observed in South China. Snowfall events are most frequent

523 in Northeast China, while are most intense in Central China. Conductor galloping

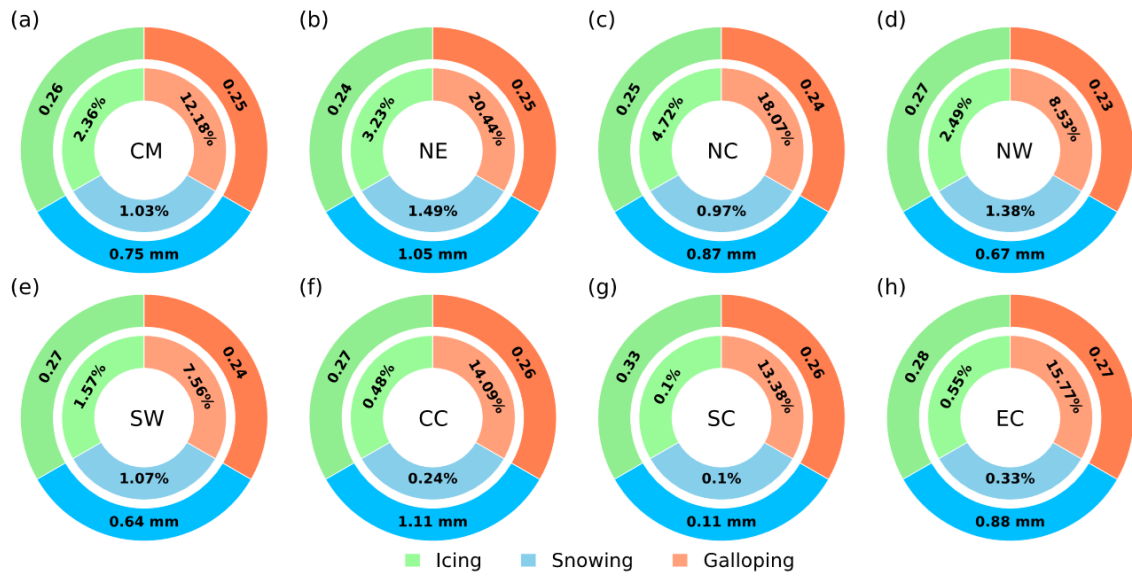
524 events are most common in Northeast China while their peak intensity is found in East

525 China. On average, the frequency and mean intensity of ice accretion, snowfall and  
 526 conductor galloping events in China are 2.36% and 0.26, 1.03% and 0.75 mm, and  
 527 12.18% and 0.25, respectively. In the demand-side, both the frequency and intensity of  
 528 extreme high temperature are relatively high in Northwest and South China. Extreme  
 529 low temperature reach its highest frequency and intensity in Northeast China. Similarly,  
 530 heat and humid environment is most pronounced in South, East and Central China. On  
 531 average, the frequency and mean intensity of extreme high temperature, extreme low  
 532 temperature and heat and humid environment in China are 0.73% and 40.94 °C, 24.84%  
 533 and -15.06 °C, and 6.07% and 0.24, respectively.



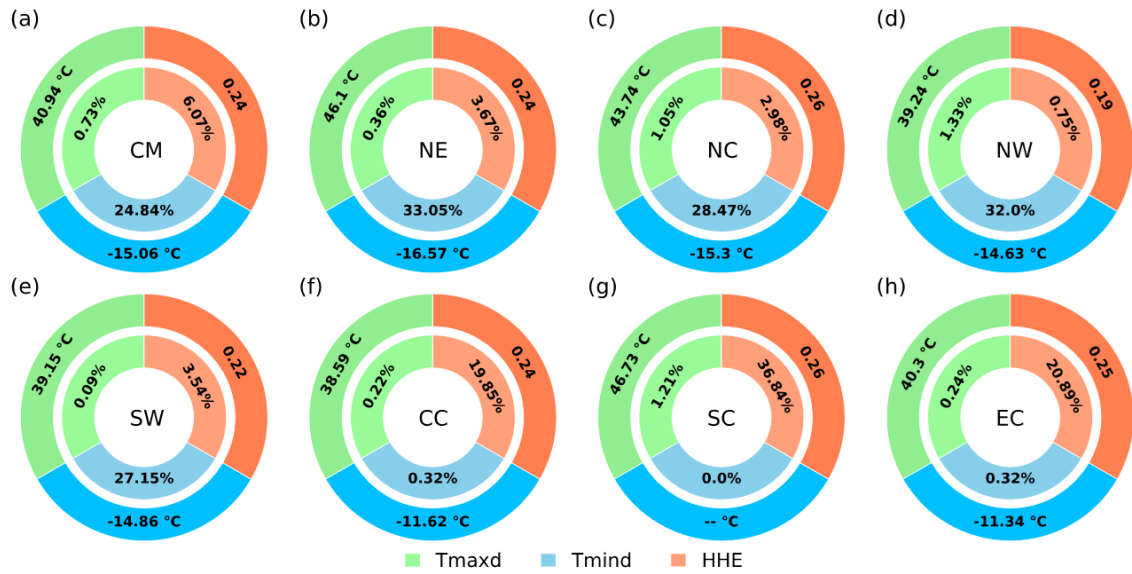
534

535 Figure 9. The annual mean frequency (unit: %/a) and intensity of high-impact weather events  
 536 relevant to generation-side across different regions of China (1980 to 2016). The inner and outer  
 537 circles correspond to the frequency and average intensity, respectively.



538

539 Figure 10. Similar to Fig. 9, but for grid-side. Note that the intensity of ice accretion and conductor  
 540 galloping events is calculated based on *CWI* indice, which is dimensionless.



541

542 Figure 11. Similar to Fig. 9, but for demand-side. Note that the intensity of heat and humid  
 543 environment events is calculated based on *CWI* indice, which is dimensionless.

544 **5. Concluding remarks**

545 In new-type power systems dominated by wind and solar energy, there is a

546 pronounced “weather dependency” and “system vulnerability”, where high-impact  
547 weather events can amplify risks to operational security. Developing a high-quality  
548 gridded dataset that involves both meteorological variables and high-impact weather  
549 events is of great significance. In this study, the China New-type Power Systems  
550 Meteorological (CNPS-Met) dataset is developed, and the spatiotemporal  
551 characteristics of high-impact weather events affecting new-type power systems are  
552 analyzed. The main conclusions are summarized as follows:

553         An improved optimal interpolation assimilation scheme, herein referred to as the  
554 spatially adaptive optimal interpolation scheme, is employed to generate the CNPS-Met  
555 dataset. Unlike conventional optimal interpolation schemes that utilize a fixed influence  
556 radius in the observation operator, the improved scheme adaptively adjusts the  
557 influence radius based on the spatial density and distribution of observational stations,  
558 thereby providing the capability to effectively characterize local variations in  
559 meteorological variables.

560         The CNPS-Met dataset covers the entire Chinese mainland. It features a daily  
561 temporal resolution and a 25 km spatial resolution. The dataset includes eight  
562 meteorological variables and eleven high-impact weather events, categorized into  
563 generation-side, grid-side and demand-side perspectives. Evaluation results indicates  
564 that, the meteorological estimates from the CNPS-Met dataset generally demonstrate  
565 superior performance compared to the other three datasets (CN05.1, CMFD and  
566 CDMet). This advantage is consistent across various meteorological variables and  
567 throughout most regions of China, as evidenced by lower *MREs* and higher *GPI* values.

568 Furthermore, CNPS-Met maintains higher accuracy in most years, seasons, and months.  
569 Compared to the other datasets, the estimated *MREs* of 2-m air temperature, 2-m  
570 specific humidity, 10-m wind speed, precipitation and surface pressure averaged over  
571 the Chinese mainland from 1980 to 2016 in CNPS-Met could be reduced by 1.7%-  
572 18.5%, 9.0%-29.6%, 1.9%-8.5%, 2.7%-18% and 4.9%-5.2%, respectively.

573 Based on the observation experiments, ideal experiments, and literature research,  
574 a series of high-impact weather events critical to the operation of new-type power  
575 systems are identified. In the generation-side, the frequency and mean intensity of cut-  
576 out wind speed, cut-in wind speed, low radiation, extreme high temperature and  
577 extreme low temperature in China are 0.4% and  $37.3 \text{ m s}^{-1}$ , 58.9% and  $1.5 \text{ m s}^{-1}$ , 14.9%  
578 and  $30.1 \text{ W m}^{-2}$ , 2.5% and  $37.1 \text{ }^{\circ}\text{C}$ , 9.9% and  $-23.1 \text{ }^{\circ}\text{C}$ , respectively. In the grid-side,  
579 the frequency and mean intensity of ice accretion, snowfall and conductor galloping  
580 events in China are 2.36% and 0.26, 1.03% and 0.75 mm, and 12.18% and 0.25,  
581 respectively. In the demand-side, the frequency and mean intensity of extreme high  
582 temperature, extreme low temperature and heat and humid environment in China are  
583 0.73% and  $40.94^{\circ}\text{C}$ , 24.84% and  $-15.06 \text{ }^{\circ}\text{C}$ , and 6.07% and 0.24, respectively.

584 Results of this study are anticipated to establish a foundation for research and  
585 applications spanning meteorology and new-type power systems, and are expected to  
586 ultimately support the formulation of renewable energy policies in China. Our future  
587 work will focus on investigating the direct (e.g., damage to, failure of, and performance  
588 degradation in power generation equipment) and indirect (e.g., reduced power  
589 generation efficiency and increased operation and maintenance costs) impacts of

590 meteorological conditions on the generation-side, grid-side, and demand-side of the  
591 new-type power system through field observations or idealized experiments, thereby  
592 establishing a more comprehensive and scientific identification for high-impact weather  
593 events, especially the compound weather events. Additionally, influences of high-  
594 impact weather events on wind and solar energy are different, which will also be  
595 investigated. Furthermore, our dataset is designed to be a living dataset that can be  
596 continuously extended, we shall update this dataset continuously and enhance the  
597 spatiotemporal resolution and quality of the CNPS-Met dataset by applying artificial  
598 intelligence methods (including image enhancement techniques etc.) and incorporating  
599 underlying surface characteristics and satellite data.

600 A detailed description of the CNPS-Met dataset is provided in Table 2.

601 Table 2. Introduction to the CNPS-Met dataset.

Entry	Descriptions
Spatial coverage	The Chinese Mainland (excluding maritime territorial)
Temporal range	1980-current (ongoing updates)
Spatial resolution	25 km×25 km
Temporal resolution	Daily
Time Standard	Universal Time Coordinated (UTC)
Format	NetCDF
Invalid value	-999.0
Abbreviation and introduction of meteorological variables	tas: 2-m mean temperature; tmax: 2-m maximum temperature; tmin: 2-m minimum temperature; precip: accumulated precipitation; wind: 10-m mean wind speed; rhum: 2-m mean relative humidity; shum: 2-m mean specific humidity; pres: mean surface pressure
Abbreviation for high-impact weather events in three critical vulnerability dimensions	Generation-side: Vout, Vin, Lowrad, Tmaxg, Tming Grid-side: Icing, Snowing, Galloping Demand-side: Tmaxd, Tmind, HHE

602 The file name for CNPS-Met follows the pattern:

603 CNPS\_Type\_History\_Daily\_Variable\_CCYY.nc, and all times are in Coordinated  
604 Universal Time (UTC). In this naming convention: “Type” is an abbreviation for  
605 meteorological variables and for the generation side, grid side, and demand side of the  
606 new power system, represented respectively by “Meteo”, “Generation”, “Grid”, and  
607 “Demand”, respectively; “Variable” is an abbreviation for the variable name; “CCYY ”  
608 represents the year (e.g., 1980,1981, .....).

609 The meteorological variables include: tas (2-m mean temperature), tmax (2-m  
610 maximum temperature), tmin (2-m minimum temperature), precip (accumulated  
611 precipitation), wind (10-m mean wind speed), rhum (2-m mean relative humidity),  
612 shum (2-m mean specific humidity), pres (mean surface pressure). The high-impact  
613 weather on the generation side includes: Vout (cut-out wind speed), Vin (cut-in wind  
614 speed), Lowrad (low radiation), Tmaxg (extreme high temperature), Tming (extreme  
615 low temperature). The high-impact weather on the grid-side includes: Icing (ice  
616 accretion), Snowing (snowfall), Galloping (conductor galloping). The high-impact  
617 weather on the demand-side includes Tmaxd (extreme high temperature), Tmind  
618 (extreme low temperature), and HHE (heat and humid environment).

619

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623

624

625 **Data availability statement**

626 The CNPS-Met dataset is available in its most updated version from our public  
627 repository at <https://www.doi.org/10.12072/ncdc.nieer.db6972.2025> (Zhang et al.  
628 2025). Data are provided as standard NetCDF format. Unit conventions and detailed  
629 variable descriptions are included in the metadata and the paper.

630

631 **Author contributions**

632 FZ: data curation, conceptualization, methodology, writing–original draft,  
633 writing–review and editing. KB: methodology, data analysis and visualization, writing–  
634 review and editing. XC: project administration, funding acquisition, writing–review  
635 and editing. YY: supervision, writing–review and editing, project. FY: project  
636 administration, funding acquisition. CW: supervision, conceptualization, writing–  
637 review and editing.

638

639 **Competing interests**

640 The contact author has declared that none of the authors has any competing  
641 interests.

642

643 **References**

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