



# 1 Global Ocean Data Set of Marine Aerosol Properties

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11 **Abstract.** NOAA's Pacific Marine Environmental Laboratory (PMEL) has made measurements of  
12 aerosol chemical, microphysical, optical, and cloud nucleating properties onboard research cruises since  
13 1991. The twenty-five cruises have covered all of the world's oceans -- the Pacific, Atlantic, Indian,  
14 Arctic, and Southern. The result is the most comprehensive, publicly available database to date of aerosol  
15 properties in the marine atmosphere. The database also contains gas phase species (O<sub>3</sub>, SO<sub>2</sub>, Radon, and  
16 dimethylsulfide (DMS), seawater species (DMS, NH<sub>4</sub><sup>+</sup>, NO<sub>3</sub><sup>-</sup>, and chlorophyll-a), and meteorological  
17 parameters. Details of the cruises (locations, dates, and objectives), parameters measured, instrumentation  
18 used, and data availability are provided here. Also included are PMEL's high-level major findings and  
19 past usage of the data by others. The goal of this paper is to promote broader awareness of the database  
20 to the atmospheric aerosol *in situ* measurement, satellite, and modelling communities. Data are publicly  
21 available at NOAA's National Centers for Environmental Information (NCEI) data archive  
22 (<https://www.ncei.noaa.gov/>). Links to the Digital Object Identifiers (DOIs) for each cruise are provided  
23 herein.

## 24 25 1 Introduction

26  
27 Aerosol particles influence Earth's radiation budget directly by scattering and absorbing incoming solar  
28 radiation and indirectly by acting as cloud condensation nuclei (CCN) and impacting cloud properties  
29 including reflectivity, lifetime, and spatial extent. The concentration and composition of aerosol particles  
30 vary both geographically and temporally, leading to a wide range of local and global climate impacts.  
31 Over oceans, aerosol particles have both continental and oceanic sources. Particles emitted from  
32 continental sources, including fossil fuel combustion, biomass burning, dust, and biogenic emissions, can  
33 be transported hundreds to thousands of kilometers over oceans either in the boundary layer or the free



34 troposphere (FT). Aerosol number and mass concentrations, chemical composition, and optical and cloud-  
35 nucleating properties are impacted by transport events and vary with distance from shore (Quinn et al.,  
36 2015). The ocean itself is a source of aerosol particles through wave-breaking at the surface and  
37 subsequent bubble bursting. In addition, marine vessel emissions also contribute to the aerosol population  
38 over oceans, particularly in coastal regions and major shipping lanes (Corbett et al., 2007). Marine aerosol  
39 is defined here as particles in the atmosphere over oceans regardless of origin.

40

41 Observations of aerosol properties in the marine atmosphere are required to improve the accuracy of  
42 model simulations of their radiative effects. Satellite observations provide broad spatial and temporal  
43 coverage of the atmospheric aerosol burden over the world's oceans and reveal information about  
44 seasonally persistent transport from continents. Examples include the transport of African dust westward  
45 across the Atlantic every summer (Kaufman et al., 2005) and transport of Asian dust and pollution  
46 eastward across the Pacific during the spring (Logan et al., 2010). While satellite observations have the  
47 advantage of providing global coverage, *in situ* observations have the highest level of accuracy available  
48 to constrain radiative forcing and reduce uncertainties in forcing estimates (Li et al., 2022). As such, *in*  
49 *situ* measurements provide detailed information about the processes controlling variability in aerosol  
50 properties due to complex particle and gas phase precursor sources, transport pathways, and removal  
51 processes.

52

53 Cruises conducted since 1991 by PMEL cover all of the world's ocean providing the most  
54 comprehensive, publicly available global database to date of marine aerosol microphysical, chemical,  
55 optical, and cloud-nucleating properties. The cruises were process oriented geared toward understanding  
56 the effects of formation, emission, atmospheric transformation and removal on aerosol properties. Some  
57 of the cruises were conducted during the time of the year when a targeted plume was expected to be most  
58 pronounced. For those cruises, the range of reported values most likely is skewed toward higher values  
59 that are typical of seasonally maximum plumes. In addition, reported variability in the data are based on  
60 a snapshot during the short-lived campaigns. Objectives are described and references are cited to provide



61 context for each cruise. Data were collected using standardized methods and sampling protocols to  
62 eliminate biases in the data and to allow for direct comparison between cruises.

63

64 This paper describes the measurements in detail, the data that are available for each cruise, PMEL's major  
65 findings, and data usage by others. The goal is to provide information about data availability and to  
66 advance the widespread use of the data to the atmospheric aerosol *in situ* measurement, satellite, and  
67 modelling communities. Sect. 2 describes the cruises and Sect. 3 describes the methods. PMEL's major  
68 findings are summarized in Sect. 4., data usage by others in Sect. 5, and a brief summary in Sect. 6. Data  
69 availability is described in Sect. 7.

70

## 71 **2 Global Ocean Cruises**

72

73 Ship tracks of PMEL's cruises between 1991 and 2020 are shown in Figure 1. A list of the cruises with  
74 start and stop dates, departure and arrival ports, location, and relevant references is provided in Table 1.  
75 A complete list of instrumentation on each cruise is presented in Section 3. The Pacific Stratus Sulfur  
76 Investigation, PSI-91 is the first cruise reported here. It took place in spring of 1991 in the eastern North  
77 Pacific off the coast of Washington state with the NOAA *R/V Discoverer* leaving from Seattle, WA in  
78 mid-April and returning in early May. Measurements focused on the role of DMS oxidation products in  
79 new particle production versus particle growth (Quinn et al., 1993; Covert et al., 1992) and the seawater  
80 sulfur cycle (Bates et al., 1994).

81

82 The Marine Aerosol and Gas Exchange cruise (MAGE92) took place in the tropical Pacific in 1992 with  
83 the USC *R/V John Vickers* leaving from Los Angeles, CA in mid-February, transiting to Nuka Hiva in  
84 the Marquesas Islands, and then returning to Los Angeles in late March. Similar to PSI-91, the goals of  
85 MAGE92 were to assess the seawater sulfur cycle and processes controlling the atmospheric aerosol  
86 particle number size distribution in the marine boundary layer (MBL) (Covert et al., 1996). In addition,  
87 instrumentation was augmented to include an integrating nephelometer to measure the aerosol light  
88 scattering coefficient at 550 nm (Charlson et al., 1967). The measurements were used to assess variability

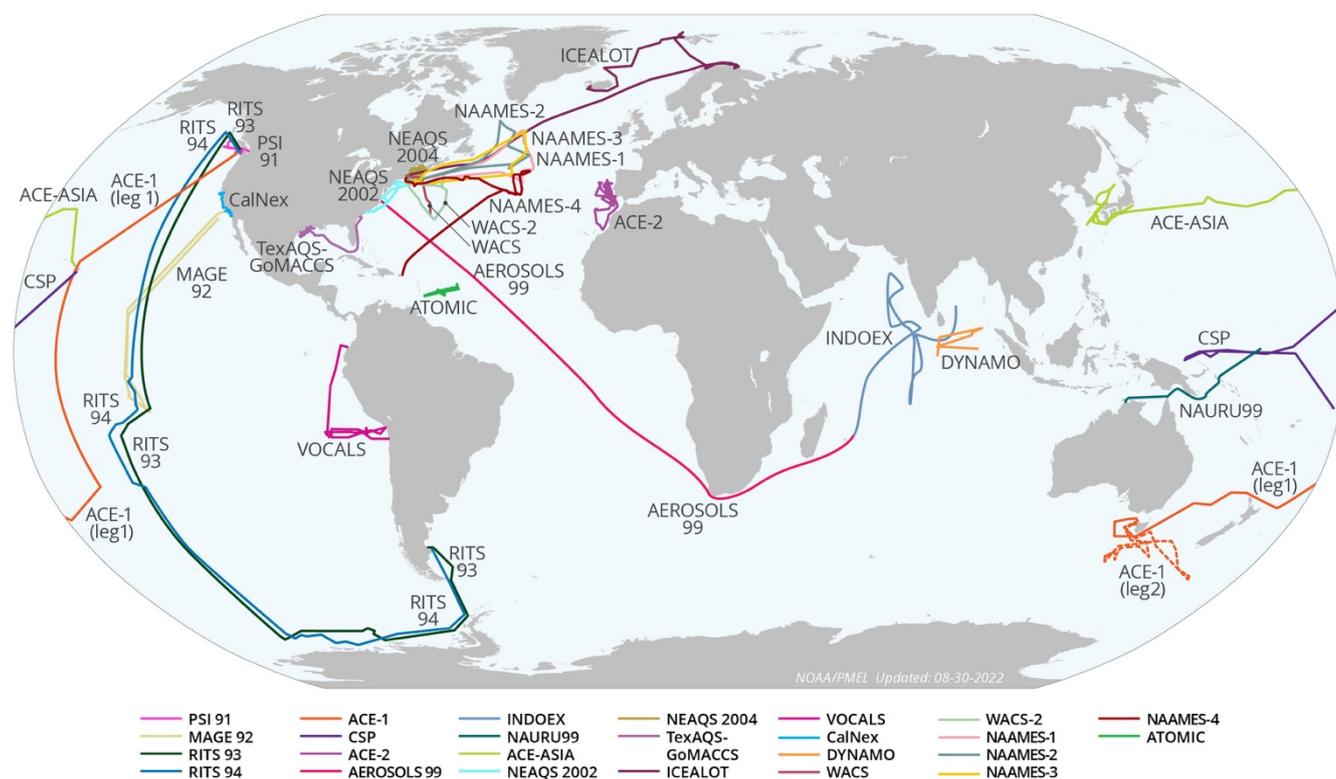


89 in aerosol chemical, microphysical, and optical properties relevant to direct radiative forcing (Quinn et  
90 al., 1995).

91

92 **Figure 1. Cruise tracks for PMEL cruises between 1991 and 2020.**

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95

96 The Radiatively Important Trace Species (RITS) cruises, RITS93 and RITS94, took place between March  
97 and May 1993 and November 1993 and January 1994, respectively. During RITS93, the NOAA *R/V*  
98 *Surveyor* went from Palmer Station, Antarctica to the Gulf of Alaska while RITS94 went in the opposite  
99 direction. These cruises extended the measurements made during MAGE92 to the central Pacific between  
100 55°N and 70°S (Quinn et al., 1996; Covert et al., 1996; Anderson et al., 1996).

101

102 In 1995, a series of Aerosol Characterization Experiments (ACE) was initiated under the auspices of the  
103 International Global Atmospheric Chemistry (IGAC) Project. The overall goal of ACE was to quantify



104 the chemical and physical processes controlling the properties and evolution of aerosol particles relevant  
105 to radiative forcing and climate. Each experiment was multi-platform (research ships, aircraft, and ground  
106 stations) with international participation. The first Aerosol Characterization Experiment (ACE-1) took  
107 place in the Southern Ocean to target aerosol in a remote, minimally polluted marine atmosphere as a  
108 reference for future experiments (Bates et al., 1998a). PMEL conducted measurements onboard the  
109 NOAA *R/V Discoverer* during Leg 1 from Seattle, WA, USA to Hobart, Australia and then Leg 2 in the  
110 Southern Ocean in and out of Hobart, Australia. ACE-1 extended earlier Pacific measurements to the  
111 Southern Ocean and augmented the characterization of optical properties through the addition of a single  
112 wavelength (550 nm) Particle Soot Absorption Photometer (PSAP) (Quinn et al., 1998b).

113

114 Subsequent ACEs were conducted downwind of continents to characterize changes in aerosol properties  
115 with advection over the ocean. ACE-2 was conducted in June and July of 1997 over the sub-tropical  
116 northeast Atlantic to characterize pollution and dust aerosol as it was advected from Europe and Africa  
117 and mixed into the marine atmosphere (Raes et al., 2000). PMEL made measurements onboard the  
118 Institute of Biology of the Southern Seas (IBSS) *R/V Professor Vodyanitskiy* leaving from and returning  
119 to Lisbon, Portugal. Background marine, anthropogenic, and dust aerosol were encountered (Bates et al.,  
120 2000). Instrumentation was unchanged from that of previous cruises.

121

122 The third experiment in the series was the Indian Ocean Experiment (INDOEX) which targeted the Indo-  
123 Asian haze during the Northern Hemisphere dry monsoon as it was advected over the Indian Ocean  
124 (Ramanathan et al., 2001). PMEL participated onboard the NOAA *R/V Ronald H. Brown* in a leg which  
125 brought the ship from Norfolk, VA to Mauritius during January and February of 1999. This leg was named  
126 AEROSOLS99. The second leg, officially INDOEX, took the ship from Mauritius northeast throughout  
127 the South Atlantic and Indian Oceans and ended in the Maldives during February and March of 1999.  
128 During both legs, marine background, anthropogenic, dust, and biomass burning aerosol were measured  
129 (Quinn et al., 2001). Trace element concentrations able to identify and quantify dust were added to the  
130 PMEL instrument payload for AEROSOLS99 and INDOEX.



131 **Table 1. PMEL's cruises between 1991 and 2020 with start and stop dates, departure and arrival**  
 132 **ports, ocean, and relevant references.**  
 133

	Dates		Ports		Ocean(s)	References
	Start	Stop	Departure	Arrival		
PSI-91 <sup>a</sup>	4/15/1991	5/1/1991	Seattle, WA, USA	Seattle, WA, USA	North Pacific (coastal Washington)	Covert et al. (1992); Quinn et al. (1993)
MAGE92 <sup>b</sup>	2/21/1992	3/25/1992	Los Angeles, CA, USA	Nuka Hiva, Marquesas Islands	Tropical Pacific	Quinn et al. (1995)
RITS93 <sup>c</sup>	3/20/1993	5/7/1993	Punta Arenas, Chile	Seattle, WA, USA	South and Tropical Pacific	Covert et al. (1996)
RITS94 <sup>d</sup>	11/20/1993	1/7/1994	Seattle, WA, USA	Punta Arenas, Chile	North and Tropical Pacific	Covert et al. (1996)
ACE-1 <sup>e</sup> Leg 1	10/12/1995	11/9/1995	Seattle, WA, USA	Hobart, Australia	Pacific	Bates et al. (1998a)
ACE-1 <sup>e</sup> Leg 2	11/15/1995	12/13/1995	Hobart, Australia	Hobart, Australia	Southern Ocean	Bates et al. (1998a)
CSP <sup>f</sup>	3/12/1996	4/13/1996	Pago Pago, American Samoa	Honolulu, HI, USA	Tropical Pacific	Post et al. (1997)
ACE-2 <sup>g</sup>	6/18/1997	7/24/1997	Lisbon, Portugal	Lisbon, Portugal	Northeast Atlantic	Raes et al. (2000)
AEROSOLS99	1/14/1999	2/8/1999	Norfolk, VA, USA	Cape Town, South Africa	Atlantic	Bates et al. (2001)
INDOEX <sup>h</sup>	2/22/1999	3/30/1999	Mauritius	Male, Maldives	South Atlantic and Indian	Ramanathan et al. (2001)
NAURU99 <sup>i</sup>	6/15/1999	7/19/1999	Darwin, Australia	Kwajalein, Marshall Islands	Tropical Pacific	Post et al. (2000)
ACE-Asia <sup>j</sup>	3/15/2001	4/20/2001	Honolulu, HI, USA	Yokosuka, Japan	Western Pacific	Bates et al. (2004); Huebert et al. (2003)
NEAQS 2002 <sup>k</sup>	7/12/2002	8/11/2002	Charleston, SC, USA	Charleston, SC, USA	Gulf of Maine, Atlantic Ocean	Bates et al. (2005)
NEAQS 2004 <sup>l</sup>	7/5/2004	8/12/2004	Portsmouth, NH, USA	Portsmouth, NH, USA	Gulf of Maine, Atlantic Ocean	Fehsenfeld et al. (2006)
TexAQS-GoMACCS <sup>m</sup>	7/27/2006	9/11/2006	Charleston, SC, USA	Galveston, TX, USA	Gulf of Mexico	Parrish et al. (2009); Bates et al. (2008)
ICEALOT <sup>n</sup>	3/19/2008	4/24/2008	Woods Hole, MA, USA	Reykjavik, Iceland	North Atlantic, Arctic Ocean	Quinn et al. (2017); Russell et al. (2010)
VOCALS <sup>o</sup>	10/13/2008	12/2/2008	Panama City, Panama	Arica, Chile	Tropical Pacific	Hawkins et al. (2010); Wood et al. (2011)
CalNex <sup>p</sup>	5/14/2010	6/8/2010	San Diego, CA, USA	San Francisco, CA, USA	California Coast	Ryerson et al. (2013); Bates et al. (2012)
DYNAMO <sup>q</sup>	9/29/2011	12/8/2011	Phuket, Thailand	Phuket, Thailand	Indian Ocean	Dewitt et al. (2013)



WACS <sup>f</sup>	8/19/2012	8/27/2012	Boston, MA, USA	St. George's, Bermuda	North Atlantic	Quinn et al. (2014); Keene et al. (2017)
WACS2 <sup>g</sup>	5/20/2014	6/5/2014	Woods Hole, MA, USA	Woods Hole, MA, USA	North Atlantic	Aller et al. (2017)
NAAMES1 <sup>h</sup>	11/6/2015	12/1/2015	Woods Hole, MA, USA	Woods Hole, MA, USA	North Atlantic	Quinn et al. (2019); Quinn et al. (2017); Behrenfeld et al. (2019)
NAAMES2 <sup>h</sup>	5/11/2016	6/5/2016	Woods Hole, MA, USA	Woods Hole, MA, USA	North Atlantic	Quinn et al. (2019); Quinn et al. (2017); Behrenfeld et al. (2019)
NAAMES3 <sup>h</sup>	8/30/2017	9/24/2017	Woods Hole, MA, USA	Woods Hole, MA, USA	North Atlantic	Quinn et al. (2019); Quinn et al. (2017); Behrenfeld et al. (2019)
NAAMES4 <sup>h</sup>	3/20/2018	4/13/2018	San Juan, Puerto Rico	Woods Hole, MA, USA	Tropical and North Atlantic	Quinn et al. (2019) Behrenfeld et al. (2019)
ATOMIC <sup>u</sup>	1/7/2020	2/13/2020	Bridgetown, Barbados	Bridgetown, Barbados	Tropical Atlantic	Quinn et al. (2021)

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 135 <sup>a</sup>Pacific Stratus Sulfur Investigation 1991  
 136 <sup>b</sup>Marine Aerosol and Gas Exchange 1992  
 137 <sup>c</sup>Radiatively Important Trace Species 1993  
 138 <sup>d</sup>Radiatively Important Trace Species 1994  
 139 <sup>e</sup>Aerosol Characterization Experiment-1 (<https://data.eol.ucar.edu/project/ACE-1>)  
 140 <sup>f</sup>Combined Sensor Program (<https://psl.noaa.gov/psd3/air-sea/csp/>)  
 141 <sup>g</sup>Aerosol Characterization Experiment-2  
 142 <sup>h</sup>Indian Ocean Experiment (<http://www.indoex.ucsd.edu/index.html>)  
 143 <sup>i</sup><https://psl.noaa.gov/psd3/air-sea/nauru99/>  
 144 <sup>j</sup>Aerosol Characterization Experiment-Asia ([https://www.eol.ucar.edu/field\\_projects/ace-asia](https://www.eol.ucar.edu/field_projects/ace-asia))  
 145 <sup>k</sup>New England Air Quality Study 2002 (<https://csl.noaa.gov/projects/neaqs/>)  
 146 <sup>l</sup>New England Air Quality Study and International Consortium for Atmospheric Research on Transport and Transformation  
 147 2004 (<https://csl.noaa.gov/projects/icartt/>)  
 148 <sup>m</sup>Texas Air Quality Study/Gulf of Mexico Atmospheric Composition and Climate Study (<https://csl.noaa.gov/projects/2006/>)  
 149 <sup>n</sup>International Chemistry Experiment in the Arctic Lower Troposphere  
 150 <sup>o</sup>VAMOS Ocean-Cloud-Atmosphere-Land Study ([https://www.eol.ucar.edu/field\\_projects/vocals](https://www.eol.ucar.edu/field_projects/vocals))  
 151 <sup>p</sup>California Research at the Nexus of Air Quality and Climate Change (<https://csl.noaa.gov/projects/calnex/>)  
 152 <sup>q</sup>Dynamics of the Madden-Julian Oscillation ([https://www.eol.ucar.edu/field\\_projects/dynamo](https://www.eol.ucar.edu/field_projects/dynamo))  
 153 <sup>r</sup>Western Atlantic Climate Study 2012  
 154 <sup>s</sup>Western Atlantic Climate Study 2014 ([https://saga.pmel.noaa.gov/field\\_WACS2](https://saga.pmel.noaa.gov/field_WACS2))  
 155 <sup>t</sup>The North Atlantic Aerosols and Marine Ecosystem Study-1 (<https://science.larc.nasa.gov/NAAMES/>)  
 156 <sup>u</sup>Atlantic Tradewind Ocean-Atmosphere Mesoscale Interaction Campaign (<https://psl.noaa.gov/atomic/>)  
 157  
 158



159 ACE-Asia, the fourth and final of the ACEs, was conducted in March through May of 2001 downwind of  
160 eastern Asia to target seasonal outbreaks of Asian dust associated with frontal systems moving to the east  
161 through dust-producing regions (Huebert et al., 2003). PMEL sampled onboard the NOAA *R/V Ronald*  
162 *H. Brown* from mid-March to mid-April in 2001 as the ship transited from Honolulu, HI to the western  
163 Pacific and then spent time east of Japan and in the Sea of Japan. Between Honolulu and 2,000 mile east  
164 of Japan, marine air minimally impacted by continental emissions was sampled. West onward from that  
165 point, air masses heavily influenced by Asian emissions were sampled (Bates et al., 2004). Measurements  
166 of organic carbon (OC) and elemental carbon (EC) were added to the instrument payload for ACE-Asia.  
167

168 During the ACE years, PMEL participated in two other cruises, the Combined Sensor Program (CSP) in  
169 March and April 1996 and NAURU99 in June and July of 1999. CSP took place in the central and tropical  
170 western Pacific with the NOAA *R/V Discoverer* leaving from Pago Pago, American Samoa in mid-March  
171 and arriving in Honolulu, HI in mid-April. The overarching goal of CSP was to better understand  
172 relationships between atmospheric and oceanic variables that affect radiative balance, including aerosol  
173 particles (Post et al., 1997). NAURU99 took place onboard the NOAA *R/V Ronald H. Brown* in the  
174 southwestern Pacific in the vicinity of Nauru Island in Papua New Guinea. NAURU99 had similar  
175 scientific goals as CSP and, in addition, was conducted to assess how representative measurements made  
176 on the islands of Nauru and Manus were of the surrounding ocean (Post et al., 2000). The ship left Darwin,  
177 Australia in mid-June and arrived in Kwajalein in the Marshall Islands in mid-July.

178  
179 In 2002, a series of air quality and climate field campaigns was initiated by NOAA with other agency and  
180 academic partners. These campaigns were designed to determine the atmospheric processes that control  
181 the production and distribution of air pollutants that impact air quality and climate in and downwind of  
182 several U.S. regions. These campaigns involved, to varying degrees, a combination of shipboard, aircraft,  
183 and ground-based measurements. The New England Air Quality Study in 2002 (NEAQS 2002) targeted  
184 factors controlling air quality in New England with measurements at a network of ground stations and a  
185 ship (Bates et al., 2005). The NOAA *R/V Ronald H. Brown* departed Charleston, SC in mid-July 2002



186 and transited northeast up the coast to New York City, Boston, and Acadia National Park in Maine. The  
187 ship returned to Charleston in mid-August 2022.

188

189 A second NEAQS in 2004 (NEAQS 2004) was conducted in conjunction with the joint North American  
190 and European International Consortium for Atmospheric Research on Transport and Transformation  
191 (ICARTT). The focus was on emissions from North America and their chemical transformations and  
192 removal during transport over the North Atlantic (Fehsenfeld et al., 2006). The NOAA *R/V Ronald H.*  
193 *Brown* left Portsmouth, NH in early July, made several transits along the coasts of Massachusetts, New  
194 Hampshire, and Maine, and across the Gulf of Maine toward Nova Scotia (Quinn et al., 2006). A  
195 Quadruple Aerosol Mass Spectrometer (Q-AMS) (Jayne et al., 2000) was added to the PMEL instrument  
196 payload for the measurement of nonrefractory (NR) species where NR refers to chemical components  
197 that vaporize (< 5 sec) at the vaporizer temperature of ~550°C.

198

199 The next in the series of Air Quality – Climate cruises was the Texas Air Quality – Gulf of Mexico  
200 Atmospheric Composition and Climate Study (TexAQS/GoMACCS) between July and September in  
201 2006. The goal was to assess the factors that control the formation and transport of air pollutants along  
202 the Gulf Coast of south eastern Texas and the impact the resulting species have on the radiative forcing  
203 of climate regionally and globally (Parrish et al., 2009). The NOAA *R/V Ronald H. Brown* left Charleston,  
204 SC at the end of July, headed south along the coast of Florida, transited across the Gulf of Mexico, and  
205 spent several weeks along the coast of Texas including in the Houston Ship Channel (Bates et al., 2008).  
206 The cruise ended mid-September in Galveston, TX. Measurements of the relative humidity dependence  
207 of light scattering and cloud condensation nuclei (CCN) concentrations were added to the existing PMEL  
208 instrument payload.

209

210 The final Air Quality – Climate field campaign was the 2010 California Research at the Nexus of Air  
211 Quality and Climate Change (CalNex) study. An emphasis was put on issues that are simultaneously  
212 relevant to both air pollution and climate including emission inventories, atmospheric transport and  
213 dispersion, atmospheric processing, and aerosol direct and indirect radiative effects (Ryerson et al., 2013).



214 The Woods Hole Oceanographic Institution (WHOI) *R/V Atlantis* left from San Diego, CA in mid-May  
215 2020, transited northward up the coast of California with incursions into the Ports of Los Angeles, Long  
216 Beach, San Francisco, and Oakland, and a trip up the Sacramento River (Bates et al., 2012).

217

218 In between TexAQS-GoMACCS and CalNex, PMEL participated in two other cruises. The International  
219 Chemistry Experiment in the Arctic Lower Troposphere (ICEALOT) took place as part of the 2008  
220 International Polar Year (Russell et al., 2010). The focus was on the sources, transport, and climate impact  
221 of anthropogenic aerosol and gas phase species in an ice-free region of the Arctic. The WHOI *R/V Knorr*  
222 left Woods Hole, MA mid-March, transited across the North Atlantic to the coast of Norway, then  
223 northwest to Svalbard, and southwest to Reykjavik, Iceland where the cruise concluded at the end of  
224 April.

225

226 The VAMOS Ocean-Cloud-Atmosphere-Land Study Regional Experiment (VOCALS), where VAMOS  
227 stands for Variability of the American Monsoon Systems, took place in October and November of 2008.  
228 VOCALS focused on assessing links between aerosols, clouds and precipitation and their impacts on  
229 marine stratocumulus radiative properties and couplings between the upper ocean and lower atmosphere  
230 (Wood et al., 2011). The NOAA *R/V Ronald H. Brown* left Panama City, Panama in mid-October,  
231 conducted several transits in the vicinity of 20°S from the coast to 85°W, and ended the cruise in Arica,  
232 Chile at the beginning of December (Hawkins et al., 2010).

233

234 DYNAMO, the Dynamics of the Madden-Julian Oscillation (MJO) field campaign was conducted to  
235 collect *in situ* observations to advance our understanding of MJO initiation processes and to improve  
236 MJO prediction (Yoneyama et al., 2013). PMEL's research focused on the effect of MJO-associated  
237 convection anomalies on aerosols in the marine boundary layer (Dewitt et al., 2013). The Scripps  
238 Institution of Oceanography *R/V Roger Revelle* left Phuket, Thailand at the end of September 2011 and  
239 transited to the vicinity of 0.1°N and 80.5°E where it was stationed for most of the experiment. The ship  
240 returned to Phuket on December 8.

241



242 Between 2012 and 2018, PMEL participated in a series of cruises to investigate the impacts of marine  
243 ecosystems on primary sea spray aerosol (SSA) and its cloud-nucleating properties. During each of these  
244 cruises a portion of the time was spent generating and sampling nascent primary SSA with Sea Sweep  
245 (Bates et al., 2012). The Sea Sweep data are available in the referenced data sets for WACS, WACS-2,  
246 and all four NAAMES cruises. These data are not discussed further since the emphasis here is on ambient  
247 aerosol. Data from the ambient atmospheric marine aerosol that was sampled when Sea Sweep was not  
248 in use are discussed here. The first Western Atlantic Climate Study (WACS) took place in 2012 and  
249 focused on the high-chlorophyll, biologically productive region of Georges Bank off the coast of Cape  
250 Cod and the low-chlorophyll, oligotrophic Sargasso Sea (Quinn et al., 2014; Kawamura et al., 2017). The  
251 NOAA *R/V Ronald H. Brown* left Boston, MA in mid-August, spent time at the high- and low-chlorophyll  
252 stations, and arrived at St. George's Bermuda at the end of August.

253

254 The second WACS (WACS2) took place in 2014. The WHOI *R/V Knorr* left Woods Hole, MA in mid-  
255 May, went east to 60°W and south to 33°S stopping for stations at a range of low to high biologically  
256 productive surface seawater. Atmospheric sampling took place between stations. The ship arrived back  
257 in Woods Hole in the beginning of June (Aller et al., 2017).

258

259 PMEL participated in the NASA sponsored North Atlantic Aerosols and Marine Ecosystems Study  
260 (NAAMES), a series of field campaigns conducted to assess the seasonal impact of the western subarctic  
261 North Atlantic phytoplankton bloom on aerosols and clouds (Behrenfeld et al., 2019). Four cruises,  
262 onboard the WHOI *R/V Atlantis*, took place between November 2015 and April 2018, with each cruise  
263 targeting specific seasonal phases of the annual plankton cycle (Quinn et al., 2019). The general cruise  
264 track included a transit from Woods Hole, MA to 40°N and 40°W, a northward transit with several stations  
265 to 55°N across a range of stages in each plankton seasonal cycle, followed by a return to Woods Hole.  
266 The exception was NAAMES-4, which left from San Juan, Puerto Rico and ended in Woods Hole. In  
267 seasonal but not chronological order, NAAMES-1 took place in November 2015 targeting the initiation  
268 of the phytoplankton blooming phase, NAAMES-4 in March and April 2018 targeting the accumulation  
269 phase, NAAMES-2 in May and June targeting the bloom climax, and NAAMES-3 in September 2017



270 targeting the declining phase of the bloom. To accommodate Sea Sweep sampling, atmospheric sampling  
271 of trace elements, NR chemical species, and total aerosol mass was not conducted during any of the  
272 NAAMES cruises.

273

274 The final cruise in the global data set to date is the Atlantic Tradewind Ocean-Atmosphere Mesoscale  
275 Interaction Campaign (ATOMIC) which took place in the tropical North Atlantic east of Barbados in  
276 early 2020 (Stevens et al., 2021). The NOAA *R/V Ronald H. Brown* left Bridgetown, Barbados in early  
277 January and spent time between Barbados and the Northwest Tropical Atlantic Station (NTAS) buoy 500  
278 nm to the northeast to gather information on shallow atmospheric convection, the effects of aerosols and  
279 clouds on the ocean surface energy budget, and mesoscale oceanic processes (Quinn et al., 2022).  
280 Measurements of trace element and total aerosol mass concentrations were reinstated for ATOMIC but  
281 concentrations of NR chemical species were not.

282

### 283 **3. Methods**

284

285 Sampling methods evolved between 1991 and 2020 as the number of parameters to be measured increased  
286 and the technical capabilities of instrumentation improved. In addition, instruments were added or  
287 removed from the PMEL payload depending on the goals of each cruise. Instrumentation and its evolution  
288 are described in detail below including the sampling inlet and methods for the measurement of aerosol  
289 microphysical, chemical, optical, and cloud-nucleating properties. In addition, measurement methods of  
290 gas phase species and surface seawater properties are provided. Parameters measured during each cruise  
291 are listed in Table 2 (aerosol microphysical and cloud-nucleating), Table 3 (aerosol chemical  
292 composition), Table 4 (aerosol optical), and Table 5 (gas phase and seawater species).

293

294 For all cruises, instrumentation was housed in one or more 8 ft (2.44 m) tall shipping container(s) outfitted  
295 with power, air conditioning, and, in some cases, water. Unistrut was installed on inside walls for the  
296 securing of instrument racks, drawers, shelves, etc. A railing surrounding the sampling inlet was installed  
297 on the roof of the container for mounting of meteorological and other sensors.



### 298 **3.1. Aerosol sampling inlet**

299

300 For all cruises, an aerosol sampling mast was mounted on top of an 8 ft (2.44 m) tall shipping container  
301 converted to a laboratory as described above and shown in Figure 2a. The container was mounted as far  
302 forward of the ship's stack as possible to minimize contamination. To maintain nominally isokinetic flow  
303 and minimize the loss of supermicron particles, the inlet at the top of the mast was rotated into the relative  
304 wind first manually (PSI-91 to NAURU-99) and then automatically under computer control (ACE Asia  
305 through ATOMIC). The mast angle was recorded for post-cruise data analysis. Air entered the inlet  
306 through a 5 cm diameter hole, passed through an expansion cone, and then into the 20 cm diameter  
307 sampling mast. The flow through the mast was  $1 \text{ m}^3 \text{ min}^{-1}$ . Individual 1.9 cm diameter stainless steel tubes  
308 extended into the base of the mast. These were connected to the various aerosol instruments in the  
309 laboratory container directly below the mast with carbon-embedded conductive tubing to prevent the loss  
310 of particles through static charging. Sampling of organics was added to the PMEL payload for ACE-Asia.  
311 Stainless steel tubing was added for the connections between the aerosol inlet and the instruments and  
312 impactors sampling for organic components.

313

314 During the initial cruise reported here, PSI-91, sample air from the mast was not conditioned, i.e., heated  
315 to control RH. Instead, aerosol was sampled at ambient RH ( $75 \pm 9\%$ ) although all CPCs (Condensation  
316 Particle Counters) had diffusion dryers upstream to reduce the RH of the sample air to less than 30%.  
317 Aerosol microphysical properties were measured continuously with periods of contamination, calibration,  
318 and downtime removed from the final data set. To avoid contamination by the ship's stack, samples for  
319 chemical analysis were collected only when the particle number concentration measured at the top of the  
320 mast was less than  $1000 \text{ cm}^{-3}$ , the relative wind speed was greater than  $3 \text{ m s}^{-1}$ , and the relative wind  
321 direction was forward of the ship's beam. This approach was employed during all cruises although the  
322 particle number concentration and relative wind speed and direction limits were varied depending on  
323 conditions.

324



325 After PSI-91, the last 1.5 m of the inlet were heated to establish a low reference relative humidity. Heating  
326 allows for constant instrumental size segregation in spite of variations in ambient RH and results in  
327 measurements of aerosol chemical, microphysical, and optical properties that are directly comparable.  
328 During MAGE 92, RITS 93, RITS 94, and ACE-1, sample air was heated above ambient temperatures to  
329 reach an RH of < 25%. The target RH for the sample air was increased to 55 to 60% for ACE-2 and  
330 following cruises because it is above the efflorescence humidity of most aerosol components and  
331 component mixtures (Carrico et al., 2003), which reduces particle bounce in impactors and simplifies  
332 thermodynamic equilibrium calculations of particle density and refractive index. There are a few  
333 exceptions to the target RH of 55 to 60%. The cold Arctic air temperatures during ICEALOT and the  
334 higher latitude portions of the NAAMES cruises made it difficult to obtain that RH. Instead, the aerosol  
335 was sampled at less than 25% RH during those conditions.

336

337 After RITS94 and before ACE-1, a temperature-controlled box was installed at the base of the sampling  
338 mast (Figure 2b and c). The box was heated above ambient temperatures to reduce cooling and  
339 condensation in sampling lines in the air-conditioned laboratory container and to maintain a uniform RH  
340 of the sampled air. Instrumental RH for the particle sizing systems is listed in Table 2.

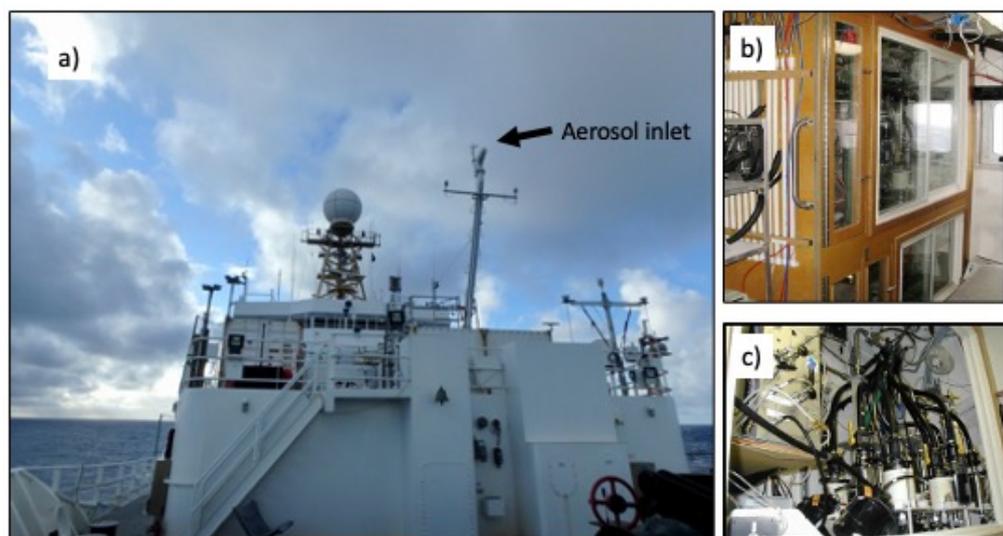
341

342 The transmission efficiency of particles through the mast as a function of size was characterized after  
343 INDOEX in the Kirsten Wind Tunnel at the University of Washington (Bates et al., 2002). The Kirsten  
344 Wind Tunnel is a subsonic, closed circuit, double return wind tunnel. The mast was mounted under the  
345 wind tunnel with the rotatable cone on top extending into the test section of the tunnel. Two sets of  
346 propellers moved air through the test section at speeds of 7 to 20 m sec<sup>-1</sup>. Aerosol particles were generated  
347 from a 10% polyethylene glycol solution (PEG-400 molecular weight mixed in distilled water) using a  
348 pressurized tank and spray nozzle downwind of the mast. Aerosol size distributions were measured from  
349 0.56 to 14 μm using Aerodynamic Particle Sizers (APS 3320, TSI, St. Paul, MN). Larger particle sizes  
350 were the focus of these tests as comparisons of total particle number concentration during ACE-1 found  
351 agreement within ~ 20% of the NCAR C-130 airplane and ground stations. The aerosol generator was

352



353 **Figure 2. Shown are a) the aerosol sampling mast mounted on top of a laboratory container onboard**  
354 **the R/V Ronald H. Brown during ATOMIC in 2020, b) the temperature-controlled box housing sizing**  
355 **instruments and impactors at the base of the mast in the laboratory container, and c) tubing connecting**  
356 **the sampling mast to sizing instrumentation (left) and impactors (right).**



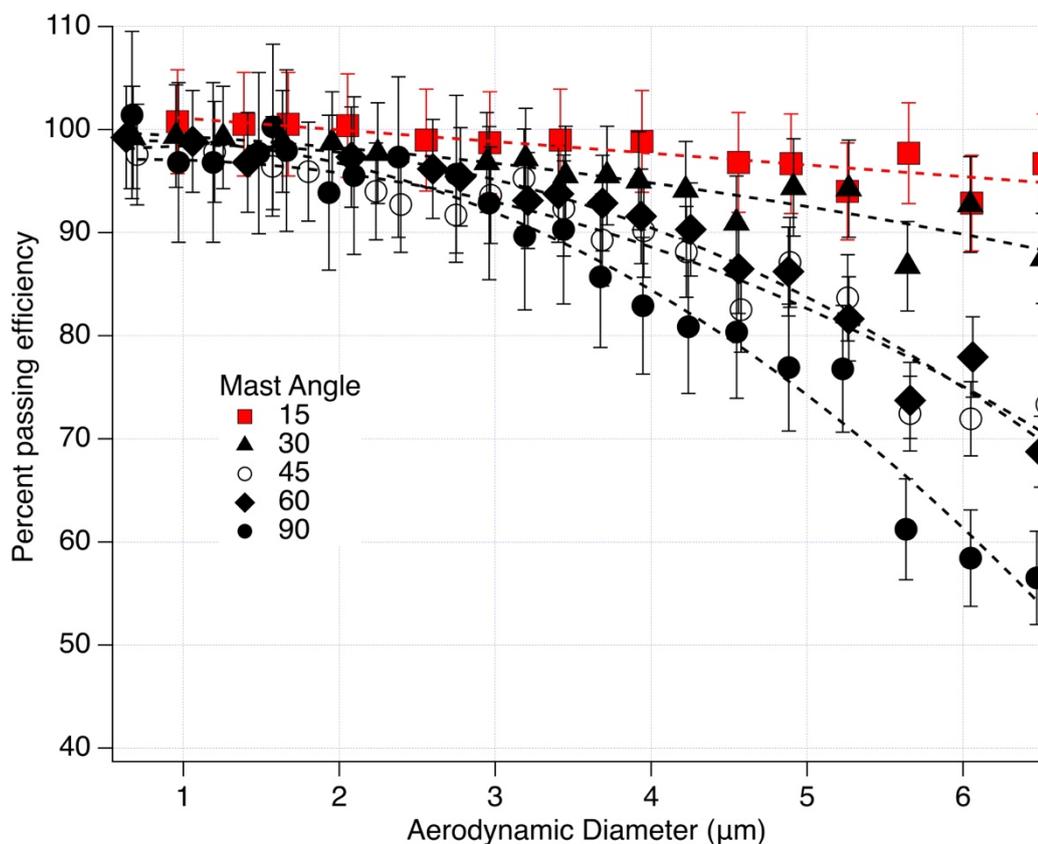
357  
358

359 operated for 1 min every 5 min to maintain a steady concentration of  $\sim 500$  particles per  $\text{cm}^3$ . Tests were  
360 conducted at different angles ( $0$  to  $90^\circ$ ) between the wind vector and the mast inlet cone axis, different  
361 wind speeds ( $7$  to  $20 \text{ m sec}^{-1}$ ), and different air flows down the mast ( $30$  to  $1200 \text{ l min}^{-1}$ ). The only  
362 parameter that was found to affect the transmission efficiency was the angle between the wind and the  
363 mast inlet cone. The transmission efficiency for particles with diameters less than  $6.5 \mu\text{m}$  was determined  
364 to be greater than  $95\%$  when the inlet was kept to within  $15^\circ$  of the wind direction (Figure 3). At a  $90^\circ$   
365 angle, the inlet transmitted about  $60\%$  of the particles in the  $6 \mu\text{m}$  size bin. Data collected in bins greater  
366 than  $6.5 \mu\text{m}$  were within the instrument noise level due to Poisson counting statistics.

367  
368



369 **Figure 3. Percent transmission efficiency of the mast at different angles between the inlet nozzle and**  
370 **the wind direction. The vertical bars indicate one standard deviation of the mean efficiency in each**  
371 **size bin. The curves are a second-order polynomial fit through the data at each angle. From Bates et**  
372 **al. (2002).**



373  
374  
375

### 376 3.2. Aerosol microphysical properties

377

378 Total particle number concentration was measured on all cruises. After the first cruise, PSI-91, number  
379 size distributions were measured on all cruises. Measurements of cloud condensation nuclei (CCN)  
380 concentrations were first added in 2006 for TexAQS. Details of the measurements are outlined below.  
381 Table 2 indicates the measurement methods used on each cruise and, in the case of number size  
382 distributions, the instrumental RH.



383

### 384 **3.2.1. Particle number concentrations**

385 As indicated in Table 2, during the first several cruises (PSI-91, MAGE92, RITS93, RITS94), total  
386 particle number concentrations for  $D_p$  greater than 3 and 12 nm were measured with TSI 3025 and 3760  
387 CPCs, respectively. The ultrafine particle number concentration was then defined as the difference  
388 between the number concentration measured with the 3025 and 3760 CPCs. For these initial cruises and  
389 all subsequent ones, diffusion driers (Permapure Inc.) were placed upstream of the CPCs to minimize  
390 particle diameter changes due to hygroscopic growth to less than 5% (Swietlicki et al., 2008). The use of  
391 a diffusion drier also helped prevent the uptake of water in the CPC condenser that results when sampling  
392 in the humid marine atmosphere. Beginning with ACE-2 in 1997 and continuing through ATOMIC in  
393 2020, particle number concentrations for  $D_p > 12$  nm were measured with a TSI 3010 CPC. Starting in  
394 2008 for VOCALS, a water-based TSI 3785 CPC was added to also measure the concentrations of  
395 particles with diameters greater than 3 nm.

396

### 397 **3.2.2. Particle number size distributions**

398 For MAGE92, RITS93, and RITS94, particle number size distributions from 0.02 to 0.6  $\mu\text{m}$  were  
399 measured with a TSI 3071 Differential Mobility Analyzer (DMA) (Quinn et al., 1998b) with the number  
400 concentration in each bin measured with a TSI 3760 CPC. The resulting number mobility distributions  
401 were inverted to a number size distribution by using the manufacturer-provided algorithm (Keady et al.,  
402 1983) and assuming that a Fuchs-Boltzman equilibrium charge distribution resulted from a  $\text{Kr}^{85}$  charge  
403 neutralizer (TSI model 3077) on the inlet of the DMA. The number concentration was corrected for the  
404 counting efficiency of the CPC (Zang et al., 1991) and diffusion losses in the DMA (Reineking et al.,  
405 1986). The sample air passed through a diffusion drier to reduce the RH to less than 25%.

406

407 A Vienna short column Ultrafine DMPS (UDMPS) coupled to a TSI 3025 CPC was added for ACE-1  
408 and all subsequent cruises to extend the size distribution measurements to the 0.005 to 0.02  $\mu\text{m}$  size range.  
409 For both ACE-1 and CSP, the UDMPS and the TSI 3071 DMA were located outside of the temperature-  
410 controlled box. Sheath air for both resulted in a measurement RH of less than 25% RH. As for the earlier



411 cruises, the mobility distributions were inverted to number size distributions by assuming that a Fuchs-  
412 Boltzman charge distribution resulted from the Kr<sup>85</sup> charge neutralizers on the inlet of the DMAs. The  
413 data were corrected for diffusional losses (Covert et al., 1997) and size dependent counting efficiencies  
414 (Wiedensohler et al., 1997) based on pre-ACE-1 intercalibration exercises. In addition, an APS (TSI  
415 3300) was added for ACE-1 and all subsequent cruises to measure the number size distribution between  
416 0.6 and 9.6  $\mu\text{m}$ . APS diameters measured at  $\sim 40\%$  RH were converted to geometric diameters by dividing  
417 by the square root of the particle density for sea salt ( $1.9 \text{ g cm}^{-3}$ ) and dried to 10% RH assuming a sea salt  
418 growth factor of 1.5 between 10 and 40% RH (Berg et al., 1998).

419

420 For ACE-2 and all subsequent cruises, the UDMPS and TSI DMA were put into the temperature-  
421 controlled box to maintain an instrumental RH of greater than 45% (see Table 2). The APS was also  
422 transferred to the temperature-controlled box where it measured at an RH of approximately 40%.

423

424 For AEROSOLS99 and all subsequent cruises, a Vienna medium column DMPS was used to measure  
425 particles in the 0.02 – 0.9  $\mu\text{m}$  size range instead of the TSI 3071 DMA. In addition, the APS model 3300  
426 was replaced with an APS model 3320. Duplicate Vienna medium column DMPSs were deployed for  
427 AEROSOLS99 and INDOEX. One measured at 10% RH outside of the temperature-controlled box and  
428 the other measured at 55% RH inside the box.

429

430 Although the APS was located in the temperature-controlled box and its inlet was maintained at 55% RH,  
431 internal heating of the sample flow by its sheath flow and waste heat likely reduced the measurement RH  
432 (Bates et al., 2004). For ACE-Asia and all subsequent cruises, the APS sheath flow was routed outside of  
433 the instrument to equilibrate with the air temperature in the temperature-controlled box and then  
434 reintroduced to the sheath and acceleration nozzle to lower the temperature and increase the measurement  
435 RH. Also starting with ACE-Asia, densities and the associated water masses at the instrumental RH were  
436 calculated with a thermodynamic equilibrium model (AeRho) using the measured inorganic ion  
437 composition (Quinn et al., 1998a). These calculated densities were used to convert the APS data from  
438 aerodynamic to geometric diameters for merging with the DMPS data. Due to the atmospheric dust that



439 was sampled during ACE Asia, the APS data were corrected for ultra-Stokesian conditions in the  
440 instrument jet and nonspherical shape (Wang et al., 1987; Wang et al., 2002).

441

### 442 **3.2.3. Cloud condensation nuclei concentrations**

443 A CCN counter (DMT CCN-100) was added for TexAQS and several later cruises (CalNex, NAAMES-  
444 1, NAAMES-2, NAAMES-3, NAAMES-4, and ATOMIC) (see Table 2). A CCN counter was onboard  
445 during WACS and WACS2 but it sampled nascent SSA with Sea Sweep for the majority of the time.  
446 CCN concentrations were measured at supersaturations between 0.2 and 1.0%. Details of the CCN counter  
447 can be found in Roberts et al. (2005). A multijet cascade impactor (Berner et al., 1979) with a 50%  
448 aerodynamic cut-off diameter of 1.1  $\mu\text{m}$  was upstream of the CCN counter. More details about the CCN  
449 measurements can be found in Quinn et al. (2008).

450

### 451 **3.3. Aerosol chemical composition**

452 The chemical species quantified along with sample collection and analysis methods are listed for each  
453 cruise in Table 3 and described below. Starting with PSI-91, size segregated aerosol was collected with a  
454 varying combination of two-, three-, and seven-stage multijet cascade impactors (Berner et al., 1979). All  
455 impactors, except those used for the analysis of organic carbon, had a grease cup at the inlet of the  
456 impactor that was coated with silicone grease to prevent the bounce of large particles onto the downstream  
457 stages. Two stage impactors had 50% aerodynamic cut-off diameters,  $D_{\text{aero},50}$ , of 1.1 and 10  $\mu\text{m}$ ; three  
458 stage impactors had  $D_{\text{aero},50}$ , of 0.18, 1.1, and 10  $\mu\text{m}$ ; and seven stage impactors had  $D_{\text{aero},50}$ , of 0.18, 0.31,  
459 0.55, 1.1, 2.0, 4.1, and 10  $\mu\text{m}$ . To attain these size cuts, air flow through all impactors was maintained at  
460 30 lpm. Flow through the impactors was computer-controlled so that aerosol was only collected when the  
461 relative wind speed and direction along with measured particle number concentration indicated there was  
462 no contamination from the ship's stack. Chemical analysis of the substrates included ion chromatography  
463 (inorganic ions), thermal/optical analysis (OC and EC), energy dispersive X-ray fluorescence (trace  
464 elements), and gravimetric mass (total aerosol mass). The substrates used in the impactors depended on  
465 the chemical species analyzed and are described below. Blank levels were determined by placing a  
466 substrate in the impactor with no air pulled through it. Blank concentrations were subtracted from sample



467 **Table 2. Microphysical and cloud-nucleating properties measured on each cruise and the**  
 468 **instrumentation used.**

Measured parameter and method	Cruise	469
<b>UFCN<sup>a</sup>, D<sub>p</sub> &gt; 3 nm</b> TSI 3025 CPC <sup>b</sup>	PSI-91, MAGE92, RITS93, RITS94, ACE-1, CSP, ACE-2, AEROSOLS99, INDOEX, NAURU99, ACE-Asia, NEAQS 2002, NEAQS 2004, TexAQS, ICEALOT	
<b>UFCN, D<sub>p</sub> &gt; 3 nm</b> TSI 3785 CPC	VOCALS, CalNex, DYNAMO, WACS, WACS-2, NAAMES-1, NAAMES-2, NAAMES-3, NAAMES-4, ATOMIC	
<b>CN<sup>c</sup> &gt; 12 nm</b> TSI 3760 CPC	PSI-91, MAGE92, RITS93, RITS94	
<b>CN &gt; 12 nm</b> TSI 3010 CPC	ACE-2, AEROSOLS99, INDOEX, NAURU99, ACE-Asia, NEAQS 2002, NEAQS 2004, TexAQS, ICEALOT, VOCALS, CalNex, DYNAMO, WACS, WACS-2, NAAMES-1, NAAMES-2, NAAMES-3, NAAMES-4, ATOMIC	
<b>Number size distribution</b> TSI 3071 DMA <sup>d</sup> , 0.02 – 0.6 μm	MAGE92, RITS93, RITS94 (< 25% RH)	
<b>Number size distribution</b> Vienna short column UDMPS <sup>e</sup> , 0.005 – 0.02 μm TSI 3071 DMA, 0.02 – 0.6 μm TSI 3300 APS <sup>f</sup> , 0.6 – 9.6 μm	ACE-1, CSP (< 25% RH); ACE-2 (45% RH)	
<b>Number size distribution</b> Vienna short column UDMPS, 0.005 – 0.02 μm Vienna medium column DMPS <sup>g</sup> , 0.02 – 0.9 μm TSI 3320 APS, 0.6 – 9.6 μm	AEROSOLS99, INDOEX (10 and 55% RH); NAURU99, ACE-Asia, NEAQS 2002, NEAQS 2004 (55% RH); TexAQS, VOCALS, CalNex, DYNAMO, CalNex, WACS, WACS-2, ATOMIC (60% RH); ICEALOT (< 25% RH); NAAMES-1, NAAMES-2, NAAMES-3, NAAMES-4 (<30 and 60% RH)	
<b>CCN<sup>h</sup></b> DMT CCN-100	TexAQS, ICEALOT, CalNex, NAAMES-1, NAAMES-2, NAAMES-3, NAAMES-4, ATOMIC	

470 <sup>a</sup>Ultrafine Condensation Nuclei  
 471 <sup>b</sup>Condensation Particle Counter  
 472 <sup>c</sup>Condensation Nuclei  
 473 <sup>d</sup>Differential Mobility Analyzer  
 474 <sup>e</sup> Ultrafine Differential Mobility Particle Sizer  
 475 <sup>f</sup>Aerodynamic Particle Sizer  
 476 <sup>g</sup>Differential Mobility Particle Sizer  
 477 <sup>h</sup>Cloud Condensation Nuclei  
 478  
 479  
 480



481 concentrations.

482

483 Additional chemical analyses were performed with a particle-into-liquid-sampler (PILS) followed by ion  
484 chromatography and water soluble organic carbon (WSOC) analysis and an aerosol mass spectrometer  
485 (AMS) for non-refractory (NR) analytes. Details are provided below.

486

### 487 **3.3.1. Inorganic ions**

488 A seven stage multi-jet cascade impactor was used on all cruises to collect size segregated samples for  
489 ion chromatography analysis. The ions quantified were  $\text{Na}^+$ ,  $\text{NH}_4^+$ ,  $\text{K}^+$ ,  $\text{Ca}^{2+}$ ,  $\text{Mg}^{2+}$ ,  $\text{Cl}^-$ ,  $\text{NO}_3^-$ ,  $\text{SO}_4^{2-}$ , and  
490 methane sulfonate or  $\text{MSA}^-$ . A Millipore Fluoropore filter (1.0  $\mu\text{m}$  pore size) was used for the final,  
491 smallest size range stage. The Millipore filter has a collection efficiency of 99% or greater for particles  
492 with diameters larger than 0.035  $\mu\text{m}$  (Liu et al., 1976). Tedlar films were used for the six largest stages.  
493 The films were cleaned in an ultrasonic bath in 10%  $\text{H}_2\text{O}_2$  for 30 min, rinsed 6 times in distilled, deionized  
494 water, and then dried in an  $\text{NH}_3$ - and  $\text{SO}_2$ -free glove box. Material collected on the filters and films was  
495 extracted by wetting with 1 mL of methanol and then adding 5 mL of distilled deionized water and  
496 sonicating for 30 min. Samples were handled in a glove box that was purged with air that had passed  
497 through a scrubber containing potassium carbonate, citric acid, and activated charcoal to remove  $\text{SO}_2$ ,  
498  $\text{NH}_3$ , and volatile organics, respectively. Sampling times varied between ~12 and 36 hrs and were based  
499 on the amount of aerosol present.

500

501 For the first few cruises (MAGE92, RITS93, RITS94), higher time resolution (< 12 hrs) submicron  
502 aerosol samples were collected using a filter holder downstream of a cyclone with a  $D_{\text{aero},50}$  of 1  $\mu\text{m}$ .  
503 Starting with ACE-1, a 2-stage impactor was used for higher time resolution sampling of sub- and  
504 supermicron aerosol for ion chromatography analysis. Substrates, sample handling, and blank  
505 determinations were the same as discussed above.

506

507 A Particle-Into-Liquid-Sampler (PILS) coupled to an ion chromatograph was used to sample submicron  
508 inorganic ions during NEAQS 2002, NEAQS 2004, and TexAQS at a higher time resolution (15 min)



509 than any of the impactors (Bates et al., 2008). The common aerosol inlet was used to deliver aerosol to a  
510 PILS at 55% RH. An impactor with a  $D_{aero,50}$  of 1.1  $\mu\text{m}$  was upstream of the PILS. Flow through the  
511 impactor was 30 slpm with 15 slpm through the PILS and 15 slpm through a bypass line. Two annular,  
512 glass denuders (URG) were in series downstream of the impactor and upstream of the PILS. One was  
513 coated with sodium carbonate for the removal of gas phase acids and the other was coated with citric acid  
514 to remove gas phase bases. Two Kloehn syringe pumps were used to deliver a solution of LiF to the top  
515 of the PILS impactor to correct for dilution of the sample within the PILS. Two additional pumps  
516 delivered sample from the PILS simultaneously to a cation and an anion IC. More information about the  
517 PILS can be found in (Weber et al., 2001). Between every 45 min to 2 hrs, sample air was passed through  
518 a HEPA filter for 15 min to remove particles and determine the measurement blank. This blank was  
519 subtracted from the sample concentrations.

520

521 For both the impactor and the PILS data, non-sea salt (nss)  $\text{SO}_4^-$  concentrations were calculated from  $\text{Na}^+$   
522 concentrations and the ratio of sulfate to sodium in seawater. Sea salt concentrations were calculated from

523

$$524 \quad \text{Sea salt } (\mu\text{g m}^{-3}) = \text{Cl}^- (\mu\text{g m}^{-3}) + \text{Na}^+ (\mu\text{g m}^{-3}) \times 1.47 \quad (1)$$

525

526 where 1.47 is the seawater ratio of  $(\text{Na}^+ + \text{K}^+ + \text{Mg}^{+2} + \text{Ca}^{+2} + \text{SO}_4^- + \text{HCO}_3^-)/\text{Na}^+$  (Holland, 1978). This  
527 approach prevents the inclusion of non-sea salt  $\text{K}^+$ ,  $\text{Mg}^{+2}$ ,  $\text{Ca}^{+2}$ ,  $\text{SO}_4^-$ , and  $\text{HCO}_3^-$  in the sea salt mass and  
528 allows for the loss of  $\text{Cl}^-$  mass through  $\text{Cl}^-$  depletion processes. It also assumes that all measured  $\text{Na}^+$  and  
529  $\text{Cl}^-$  is derived from seawater. Results of Savoie et al. (1980) indicate that soil dust has a minimal  
530 contribution to measured soluble sodium concentrations.

531

### 532 3.3.2. Organic and Elemental Carbon

533 Starting with ACE-Asia in 2001, sub-1 and sub-10  $\mu\text{m}$  samples were collected for OC/EC analysis using  
534 2 and 1 stage impactors, respectively (Bates et al., 2004). Each impactor had 2 quartz backup filters. OC  
535 concentrations from both impactors were corrected for blanks and artifacts using the last quartz filter in  
536 line. Aluminum foil was used as a substrate on the 1.1  $\mu\text{m}$  jet plate. All substrates, aluminum foil and



537 quartz, were prebaked at 500° prior to sampling. One sub-10  $\mu\text{m}$  and one sub-1  $\mu\text{m}$  impactor were  
538 operated without a denuder upstream to avoid losses of large particles in the denuder. OC from the sub-  
539 1  $\mu\text{m}$  impactor was subtracted from the OC from the sub-10  $\mu\text{m}$  impactor to determine supermicron OC  
540 concentrations. A second sub-1  $\mu\text{m}$  impactor was operated with a denuder upstream that contained strips  
541 of carbon-impregnated glass fiber filters to remove gas phase organics. Sub-1  $\mu\text{m}$  OC and EC were  
542 quantified on the impactor samples downstream of the denuder.

543

544 During ACE-Asia and NEAQS 2002, a 7-stage impactor was used for the sampling of OC/EC providing  
545 greater size resolution. Aluminum foil substrates were used on all stages of the impactor along with 2  
546 quartz backup filters. A 3-stage impactor ( $D_{50,\text{aero}}$  of 0.18, 1.1, and 10  $\mu\text{m}$ ) was used during WACS,  
547 WACS-2, and NAAMES-1 through NAAMES-4. These cruises focused on the composition and  
548 properties of sea spray aerosol. The 3 size cuts allowed for the separation of the sub-0.18  $\mu\text{m}$  aerosol from  
549 larger size ranges in recognition that smaller particle sizes are known to be enriched in organics through  
550 the sea spray aerosol production process (Keene et al., 2007).

551

552 OC and EC concentrations on the impactor substrates were measured with a Sunset Labs thermal/optical  
553 analyzer (Birch et al., 1996). Four temperature steps were used to achieve a final temperature of 870°C  
554 in He to drive off OC. After cooling the sample down to 550°C, a He/O<sub>2</sub> mixture was introduced and the  
555 sample was heated in four temperature steps to 910°C to drive off EC. The transmission of light through  
556 the filter was measured to separate EC from any OC that charred during the initial stages of heating. No  
557 correction was made for carbonate carbon so OC included both organic and inorganic carbon. The mass  
558 of particulate organic matter (POM) was determined by multiplying the measured organic carbon  
559 concentration in  $\mu\text{g m}^{-3}$  by a factor of 2.1 in marine regions and 1.6 elsewhere (Turpin et al., 2001).

560

561 A semi-continuous real-time Sunset Labs thermal/optical analyzer was used during NEAQS 2004 for  
562 higher time resolution measurements of OC concentrations. The OC/EC analyzer was downstream of a  
563 sub-1  $\mu\text{m}$  impactor and a denuder. The analyzer collected air on a filter for 45 or 105 minutes depending



564 on OC concentrations. At the end of the sampling time the instrument analyzed the filter using the same  
565 temperature program described above. The sampling times were not long enough to measure EC above  
566 the detection limit of  $0.35 \mu\text{g m}^{-3}$ .

567

### 568 **3.3.3. Water soluble organic carbon**

569 A PILS coupled to a Total Organic Carbon (TOC) analyzer (Sievers Model 800 Turbo) was used during  
570 TexAQS to measure water soluble organic carbon (WSOC) (Bates et al., 2008). As for the PILS-IC, the  
571 PILS-WSOC was connected to the common aerosol inlet but with a stainless steel line. An impactor with  
572 a  $D_{50,\text{aero}}$  of  $1.1 \mu\text{m}$  was upstream of the PILS to sample submicron aerosols. A denuder identical to the  
573 one used in the thermal/optical analysis was downstream of the impactor and upstream of the PILS to  
574 remove gas phase organics. Two syringe pumps (Kloehn) delivered low-TOC water to the top of the PILS  
575 impactor. Two additional pumps were used to pull sample out of the PILS and into the TOC analyzer.  
576 The sample was passed through a  $0.5 \mu\text{m}$  in-line filter before entering the TOC analyzer to measure  
577 WSOC. Between every 45 min to 2 hrs, sample air was passed through a HEPA filter for 15 min to remove  
578 particles and determine the measurement background. The measurement background was subtracted from  
579 the sample air to obtain ambient WSOC ambient atmospheric concentrations.

580

### 581 **3.3.4. Trace elements**

582 Starting with AEROSOLS99, sub-1 and sub- $10 \mu\text{m}$  samples were collected for trace element analysis  
583 using impactors with a  $D_{50,\text{aero}}$  of  $1.1$  and  $10 \mu\text{m}$ , respectively (Quinn et al., 2001). Energy Dispersive X-  
584 RAY Fluorescence (ED-XRF) was used for quantification (Buck et al., 2021). Both impactors collected  
585 aerosol on  $2.0 \mu\text{m}$  pore size PALL Teflo Membrane Disc Filters. Supermicron concentrations were  
586 determined by subtracting the sub- $1.1 \mu\text{m}$  values from the sub- $10 \mu\text{m}$  values. No corrections were made  
587 for particle size or loading. Samples for XRF analysis were collected during all subsequent cruises except  
588 for VOCALS, WACS, WACS-2, and NAAMES-1 to 4.

589

590 Concentrations of dust were calculated based on measured values of Al, Si, Ca, Fe, and Ti assuming that  
591 each element was present in the aerosol in its most common oxide form ( $\text{Al}_2\text{O}_3$ ,  $\text{SiO}_2$ ,  $\text{CaO}$ ,  $\text{K}_2\text{O}$ ,  $\text{FeO}$ ,



592 Fe<sub>2</sub>O<sub>3</sub>, TiO<sub>2</sub>) (Seinfeld, 1986)). The measured elemental mass concentration was multiplied by the  
593 appropriate molar correction factor as shown below (Malm et al., 1994)

594

$$595 \quad \text{Dust} = 2.2(\text{Al}) + 2.49(\text{Si}) + 1.63(\text{Ca}) + 2.42(\text{Fe}) + 1.94(\text{Ti}). \quad (2)$$

596

597 This equation includes a 16% correction factor to account for the presence of oxides of other elements  
598 such as K, Na, Mn, Mg, and V that are not included in the linear combination. In addition, the equation  
599 omits K from biomass burning by using Fe as a surrogate for soil K and an average K/Fe ratio of 0.6 in  
600 soil (Braaten et al., 1986). Non-crustal K was calculated using the K/Al ratio (0.31) of Asian loess (Jahn  
601 et al., 2001) which is similar to the ratio in Saharan dust (0.24) and average crustal rock (0.32) (Formenti  
602 et al., 2003). Sea salt Ca was accounted for based on the ratio of Ca to Na in seawater.

603

### 604 **3.3.5. Non-refractory species**

605 Concentrations of submicron non-refractory (NR) NH<sub>4</sub><sup>+</sup>, SO<sub>4</sub><sup>=</sup>, NO<sub>3</sub><sup>-</sup>, and particulate organic matter  
606 (POM) were measured on NEAQS 2004, TexAQS, ICEALOT, VOCALS, CalNex, DYNAMO, WACS  
607 and WACS-2 with a Quadrupole Aerosol Mass Spectrometer (Q-AMS, Aerodyne Research Inc.,  
608 Billerica, MA, USA) (Jayne et al., 2000). The NR species measured by the AMS are defined here as all  
609 the chemical components that vaporize at 550°C. The AMS was downstream of an impactor with a D<sub>50,aero</sub>  
610 of 1.1 μm. The ionization efficiency of the AMS was calibrated every few days with dry monodisperse  
611 ammonium nitrate particles. Particle losses due to transmission through the aerodynamic lens were  
612 corrected by using the DMPS and APS-measured size distributions. Particle losses due to bounce off of  
613 the impactor-vaporizer were corrected using simultaneously sampled NH<sub>4</sub><sup>+</sup> and non-sea salt SO<sub>4</sub><sup>=</sup>  
614 concentrations from either the PILS-IC or the impactors (Quinn et al., 2008; Quinn et al., 2006).

615

### 616 **3.3.6. Aerosol Mass**

617 A filter holder with a Millipore Fluoropore filter (1.0 μm pore size) collected aerosol downstream of a  
618 cyclone with a D<sub>50,aero</sub> of 1 μm during RITS93 and RITS94. Filters were taken back to PMEL for  
619 gravimetric analysis to determine total submicron aerosol mass. The filters were weighed before and after



620 sample collection with a Mettler UMT2 microbalance. The microbalance was housed in a glove box  
621 maintained at a constant RH to allow each sampled filter to come into equilibrium with the same vapor  
622 pressure of water, thus reducing experimental uncertainty due to a variable lab RH. For RITS93 and  
623 RITS94 the RH was maintained at less than 30% by circulating air through a flat baffle box containing a  
624 saturated solution of  $MgCl \cdot 6H_2O$  and then through the glove box (Young, 1967). The circulated air was  
625 cleaned by passing it through a scrubber containing activated charcoal, potassium carbonate, and citric  
626 acid. Filters were equilibrated overnight in the glove box prior to weighing. Static charging, which can  
627 result in balance instabilities, was minimized by coating the walls of the glove box with a static dissipative  
628 polymer (Tech Spray, Inc.), placing an antistatic mat on the glove box floor, and exposing the filters to a  
629  $^{210}Po$  source to dissipate any built-up charge.

630

631 For ACE-1 and the other cruises listed in Table 3, a 2-stage impactor was used to collect submicron and  
632 supermicron aerosol for gravimetric analysis. Millipore Fluoropore filter (1.0  $\mu m$  pore size) and Tedlar  
633 films were used for the collection of submicron and supermicron aerosol, respectively. The Tedlar films  
634 were cleaned as described in Section 3.3.1. prior to sample collection. Both the Millipore filters and the  
635 Tedlar films were weighed before and after sampling. Millipore filters were weighed on the Mettler  
636 UMT2 microbalance and Tedlar films were weighed on a Cahn Model 29 microbalance. Both balances  
637 were housed in the RH-controlled glove box described above. For cruises with higher sampling RHs of  
638 55 to 60% (see Section 3.1.), a saturated solution of KBr was used in the baffle box.

639

640 In addition to the 2-stage impactor used during ACE-Asia, a 7-stage impactor was used for higher size  
641 resolution total aerosol mass concentrations.

642

643 All reported mass concentrations include the water mass that is associated with the aerosol on the filter at  
644 the glove box RH.

645



646 **Table 3. Measurements of aerosol chemical composition on each cruise and the instrumentation used.**

<b>Chemical Species and Measurement Method</b>	<b>Cruise</b>	
		647
		648
Inorganic ions <sup>a</sup> – submicron Filter with cyclone upstream, IC <sup>b</sup>	MAGE92, RITS93, RITS94	649
Inorganic ions – sub- and supermicron 2-stage impactor <sup>c</sup> , IC	ACE-1, CSP, ACE-2, AEROSOLS99, INDOEX, NAURU99, ACE-Asia, NEAQS 2002, NEAQS 2004, TexAQS, ICEALOT, VOCALS, CalNex, DYNAMO, WACS, WACS-2, NAAMES-1, NAAMES-2, NAAMES-3, NAAMES-4, ATOMIC	650 651 652
Inorganic ions – 7 size ranges 7-stage impactor <sup>d</sup> , IC	All cruises	653
		654
Inorganic ions – submicron PILS <sup>e</sup> , IC	NEAQS 2002, NEAQS 2004, TexAQS	655
OC and EC <sup>f</sup> – sub- and supermicron 2-stage impactor, Thermal analysis <sup>g</sup>	ACE-Asia, NEAQS 2002, NEAQS 2004, TexAQS, ICEALOT, CalNex, ATOMIC	656
OC and EC – submicron Semi-continuous real-time OC/EC with impactor upstream <sup>h</sup> , Thermal analysis	NEAQS 2004	657 658
		659
OC and EC – 3 size ranges 3-stage impactor <sup>i</sup> , Thermal analysis	WACS, WACS-2, NAAMES – 1, NAAMES-2, NAAMES-3, NAAMES-4	660
OC and EC – 7 size ranges 7-stage impactor, Thermal analysis	ACE-Asia, NEAQS 2002	661
WSOC <sup>j</sup> – submicron PILS, TOC <sup>k</sup> analyzer	TexAQS	662 663
Trace Elements <sup>l</sup> – sub and sub-10 micron 2 impactors, XRF <sup>m</sup>	AEROSOLS99, INDOEX, ACE-Asia, NEAQS 2002, NEAQS 2004, TexAQS, ICEALOT, CalNex, DYNAMO, ATOMIC	664 665
Aerosol mass – sub- and supermicron Filter with cyclone upstream, Gravimetric analysis	RITS93, RITS94	666 667
Aerosol mass – sub- and supermicron 2-stage impactor, Gravimetric analysis	ACE-1, ACE-2, AEROSOLS99, INDOEX, ACE-Asia, NEAQS 2002, NEAQS 2004, TexAQS, ICEALOT, VOCALS, CalNex, DYNAMO, WACS, WACS-2, ATOMIC	668 669
		670
Aerosol mass – 7 size ranges 7-stage impactor, Gravimetric analysis	ACE-Asia	671
NR <sup>n</sup> SO <sub>4</sub> , NH <sub>4</sub> , NO <sub>3</sub> , POM	NEAQS 2004, TexAQS, ICEALOT, VOCALS, CalNex, DYNAMO, WACS, WACS-2	672

673

674 <sup>a</sup>Na<sup>+</sup>, NH<sub>4</sub><sup>+</sup>, K<sup>+</sup>, Ca<sup>2+</sup>, Mg<sup>2+</sup>, Cl<sup>-</sup>, NO<sub>3</sub><sup>-</sup>, SO<sub>4</sub><sup>2-</sup>, MSA<sup>-</sup>

675 <sup>b</sup>Ion Chromatography

676 <sup>c</sup>Multi-jet cascade impactor with two size cuts, D<sub>50,aero</sub> of 1.1 μm and 10 μm

677 <sup>d</sup>Multi-jet cascade impactor with seven size cuts, D<sub>50,aero</sub> of 0.18, 0.31, 0.55, 1.1, 2.0, 4.1, and 10 μm

678 <sup>e</sup>Particle-Into-Liquid-Sampler

679 <sup>f</sup>Organic and elemental carbon

680 <sup>g</sup>Sunset Labs thermal/optical analyzer

681 <sup>h</sup>Sunset Labs real-time, semi-continuous thermal/optical analyzer



682 <sup>i</sup>Multi-jet cascade impactor with three size cuts,  $D_{50,aero}$  of 0.18, 1.1, and 10  $\mu\text{m}$   
683 <sup>j</sup>Water Soluble Organic Carbon  
684 <sup>k</sup>Total Organic Carbon Sievers Model 800 Turbo analyzer  
685 <sup>l</sup>Al, Si, nss Ca, Ti, Fe; two multi-jet cascade impactors one with a size cut of  $D_{50,aero} = 1.1 \mu\text{m}$  and one with a  $D_{50,aero} = 10 \mu\text{m}$   
686 <sup>m</sup>Energy dispersive X-ray fluorescence  
687 <sup>n</sup>Non-refractory  
688

### 689 **3.4. Aerosol optical Properties**

690 The optical properties measured on each cruise and the instrumentation used are listed in Table 4. Aerosol  
691 light scattering coefficients were measure on each cruise with the exception of the first one, PSI-91.  
692 Variations included measurement at a single wavelength (550 nm) or three wavelengths (450, 550, and  
693 700 nm) and measurement of sub-10 micron aerosol or sub-1 and sub-10 micron aerosol. The relative  
694 humidity dependence of light scattering,  $f(\text{RH})$ , was measured on some of the cruises as were  
695 backscattering coefficients at 450, 550, and 700 nm.

696  
697 Aerosol absorption coefficients were measured on every cruise starting with ACE-1. Initially,  
698 measurements were made at 550 nm for sub-10 micron aerosol. These measurements were expanded to  
699 sub-1 and sub-10 micron aerosol starting with NAURU99 and 3 wavelengths (467, 530, and 660 nm)  
700 starting with NEAQS 2002.

701  
702 Aerosol optical depth (AOD) was measured on all cruises except for the first one, PSI-91, with handheld  
703 sunphotometers. In addition, the NASA AMES Airborne Tracking sunphotometer (AATS-6) (Livingston  
704 et al., 2000) was used during ACE-2.

705  
706 Details of the measurements are provided in the following sections.

#### 708 **3.4.1. Aerosol Light Scattering**

709 During RITS93 and RITS94, sub-10 micron aerosol light scattering was measured with a newly  
710 developed, highly sensitive multiwavelength integrating nephelometer (Bodhaine et al., 1991). This  
711 nephelometer, with its high sensitivity, was combined with the closed geometry of the Ahlquist et al.  
712 (1967) nephelometer to develop the TSI, Inc. model 3563 (Anderson et al., 1996) that was used during



713 the rest of the cruises reported on here. The enclosed geometry allows for the calibration of the  
714 nephelometer with gases with known scattering coefficients.

715

716 The TSI Inc. model 3563 integrating nephelometer was used for all remaining cruises to measure sub-1  
717 and sub-10 micron scattering at three wavelengths (450, 550, and 700 nm). Sub-1 and sub-10 micron  
718 backscattering were measured on cruises between ACE-1 and WACS. Two single-stage impactors, one  
719 having a  $D_{50,aero}$  of 1.1  $\mu\text{m}$  and the other of 10  $\mu\text{m}$  were placed upstream of the nephelometer. A valve  
720 automatically switched between the two impactors every 15 minutes so that sampling alternated between  
721 sub-1 micron and sub-10 micron aerosol. Scattering and backscattering by the supermicron aerosol was  
722 determined by difference.

723

724 During all cruises, the nephelometer was calibrated with  $\text{CO}_2$  and zeroed with particle-free air every 3 to  
725 4 days (Quinn et al., 1996). The resulting zero offset and span factors were applied to the data. In addition,  
726 data were corrected for angular nonidealities of the nephelometer, including truncation errors and non-  
727 Lambertian illumination using the method of Anderson et al. (1998) or one similar to it.

728

729 The RH of the air sampled by the nephelometer was nominally that of the common inlet as described in  
730 Section 3.1. Heating within the nephelometer likely led to slightly lower RH's than for the sizing  
731 instruments detailed in Table 2. For ACE-Asia and all subsequent cruises the nephelometer flow path was  
732 modified so that the sheath flow was conditioned outside of the instrument case to equilibrate with the  
733 temperature-controlled box. It was then reintroduced into the sheath and acceleration nozzle to minimize  
734 heating of the sample air and lowering of the measurement RH. In addition, an RH sensor was placed in  
735 the nephelometer sensing volume.

736

### 737 **3.4.2. Relative Humidity Dependence of Light Scattering, $f(\text{RH})$**

738 As indicated in Table 4, the relative humidity dependence of scattering,  $f(\text{RH})$ , was measured on TexAQS,  
739 ICEALOT, VOCALS, CalNex, DYNAMO, WACS, WACS-2, NAAMES-1, NAAMES-2, and  
740 ATOMIC. A humidity-controlled system measured light scattering at two different relative humidities,



741 ~20% and ~85%, with two nephelometers operated in series downstream of an impactor  
742 ( $D_{50,aero} = 1.1 \mu\text{m}$ ). The first nephelometer in line measured sample air dried with a PermaPure, multiple-  
743 tube nafion dryer (model PR-94). Downstream of this nephelometer a humidifier was used to add water  
744 vapor to the sample flow using 6 microporous Teflon tubes surrounded by a heated water-jacket.  
745 Humidity was measured using a chilled mirror dew point hygrometer downstream of the second  
746 nephelometer in line. The same calibration procedure described in Section 3.4.2. was used (Quinn et al.,  
747 2022).

748

### 749 **3.4.3. Aerosol Light Absorption**

750 Between ACE-1 and INDOEX, the aerosol light absorption coefficient of sub- $10 \mu\text{m}$  was measured with  
751 a Particle Soot Absorption Photometer (PSAP, Radiance Research) at a wavelength of 550 nm and ~55%  
752 RH. Measured values were corrected for a scattering artifact, the deposit spot size, flow rate, and the  
753 manufacturer's calibration (Bond et al., 1999). Beginning with NAURU99, the absorption coefficient was  
754 measured for sub- $1$  and sub- $10 \mu\text{m}$  aerosol with the PSAP located downstream of the same impactors as  
755 the nephelometer. For NEAQS 2002 and all subsequent cruises, a modified PSAP was used to measure  
756 light absorption at three wavelengths (467, 530, and 700 nm) close to that of the TSI nephelometer for  
757 calculation of single scattering albedo (Virkkula et al., 2005). Beginning with TexAQS and all following  
758 cruises, a PermaPure nafion dryer was placed upstream of the PSAP so that the sample air was at ~25%  
759 RH. Measurement of dry air was found to reduce instrument noise.

760

### 761 **3.4.4. Aerosol Optical Depth**

762 Handheld sunphotometers were used to measure AOD for ACE-1 and all subsequent cruises. A single  
763 wavelength (550 nm) sunphotometer was used for ACE-1. A microtops unit (Solar Light Co.) was used  
764 for all other cruises measuring at 380, 440, 500, 675, and 870 nm. Units were calibrated before each cruise  
765 by either Solar Light Co. or NASA Goddard Space Flight Center (GSFC) using a Langley plot approach  
766 (Shaw, 1983). Initially, a NASA Sensor Intercomparison and Merger for Biological and Interdisciplinary  
767 Oceanic Studies (SIMBIOS) MATLAB routine was used to convert raw signal voltages to AOD. Included  
768 in the conversion is a correction for Rayleigh scattering (Penndorf, 1957) and air mass to account for the



769 curvature of the Earth (Kasten et al., 1989). Beginning with ICEALOT in 2008, data were reduced as part  
 770 of NASA's Maritime Aerosol Network (Smirnov et al., 2009).

771

772 **Table 4. Measurements of aerosol optical properties on each cruise and the instrumentation used.**

Optical Property and Measurement Method	Cruise
Scattering (550 nm <sup>a</sup> ), Sub-10 micron Integrating nephelometer	RITS93, RITS94
Scattering (450, 550, 700 nm), Sub-1 and sub-10 micron Integrating nephelometer, TSI Model 3563	ACE-1, CSP, ACE-2, AEROSOLS99, INDOEX, NAURU99, ACE-Asia, NEAQS 2002, NEAQS 2004, TexAQS, ICEALOT, VOCALS, CalNex, DYNAMO, WACS, WACS-2, NAAMES-1, NAAMES-2, NAAMES-3, NAAMES-4, ATOMIC
Backscattering (450, 550, 700 nm), Sub-1 and sub-10 micron Integrating nephelometer, TSI Model 3563	ACE-1, CSP, ACE-2, AEROSOLS99, INDOEX, NAURU99, ACE-Asia, NEAQS 2002, NEAQS 2004, TexAQS, ICEALOT, VOCALS, CalNex, DYNAMO, WACS
Absorption (550 nm), Sub-10 micron PSAP <sup>b</sup> , Radiance Research	ACE-1, ACE-2, AEROSOLS99, INDOEX
Absorption (550 nm), Sub-1 and sub-10 micron PSAP <sup>b</sup> , Radiance Research	NAURU99, ACE-Asia
Absorption (467, 530, 660 nm), Sub-1 and sub-10 micron PSAP <sup>b</sup>	NEAQS 2002, NEAQS 2004, TexAQS, ICEALOT, VOCALS, CalNex, DYNAMO, WACS, WACS-2, NAAMES-1, NAAMES-2, NAAMES-3, NAAMES-4, ATOMIC
f(RH), Sub-1 micron scattering and backscattering (450, 550, 700 nm) (25 and 85% RH) 2 Integrating nephelometers, TSI Model 3563	TexAQS, ICEALOT, VOCALS, CalNex, DYNAMO, WACS, WACS-2, NAAMES-1, NAAMES-2, ATOMIC
AOD (391, 500 nm) Handheld sunphotometer	RITS93, RITS94
AOD (375, 500, 778, 862 nm) Handheld sunphotometer	ACE-1
AOD (380, 450, 525, 864, 1021 nm) AATS-6 <sup>c</sup> , NASA AMES	ACE-2
AOD (380, 440, 500, 675, 870 nm) Handheld sunphotometer, Solar Light Co. Microtops	AEROSOLS99, INDOEX, NAURU99, ACE-Asia, NEAQS 2002, NEAQS 2004, TexAQS, ICEALOT, VOCALS, CalNex, DYNAMO, WACS, WACS-2, ATOMIC

773

774 <sup>a</sup>nm, wavelength



775 <sup>b</sup>Particle Soot Absorption Photometer  
776 <sup>c</sup>Ames Airborne Tracking Sunphotometer  
777

### 778 **3.5. Gas phase species**

779 Gas phase species that were measured include O<sub>3</sub>, SO<sub>2</sub>, Radon, and DMS. The cruises each gas was  
780 measured on and the measurement methods used are listed in Table 5. O<sub>3</sub> and SO<sub>2</sub> were measured  
781 primarily as tracers of pollution. Radon (as Rn<sup>222</sup>) was measured as an indicator of contact of the sampled  
782 air with land (Whittlestone et al., 1998b). DMS was measured due to its link to nss SO<sub>4</sub><sup>-</sup> and MSA via  
783 oxidation in the atmosphere (e.g. Andreae et al. (1985)).  
784

#### 785 **3.5.1. Ozone**

786 O<sub>3</sub> was measured on all cruises with the exception of NAURU99 and WACS. Three different ozone UV  
787 analyzers were used over the years including a Dasibi 1008-AH UV photometer, a TECO Model 49 O<sub>3</sub>  
788 Analyzer, and a TECO Model 49C O<sub>3</sub> Analyzer (Table 5). For all cruises, a ¼” ID Teflon sample line  
789 was used to draw air from the top of the aerosol common sampling mast to the O<sub>3</sub> instrument located in  
790 the lab container at the base of the mast. The loss of O<sub>3</sub> in a Teflon sampling line is approximately 5%  
791 per 30 m indicating that losses were negligible (< 3%). At intervals of 1 to 4 days, a charcoal filter was  
792 placed in the sampling line for 1 hour to determine a zero which was subtracted from the O<sub>3</sub> signal. More  
793 details can be found in Johnson et al. (1990).  
794

#### 795 **3.5.2. Sulfur dioxide**

796 SO<sub>2</sub> was measured during ACE-Asia, NEAQS 2002, NEAQS 2004, TexAQS, ICEALOT, VOCALS, and  
797 CalNex with a Thermo Environmental Instruments Model 43C trace level pulsed fluorescence analyzer.  
798 Air was drawn through the 18 m aerosol common sampling mast at 1 m<sup>3</sup> min<sup>-1</sup>. At the base of the mast, a  
799 5.0 L min<sup>-1</sup> flow was pulled in series through a 1 m long Teflon tube, a Millipore Fluoropore Teflon filter  
800 (1.0 μm pore size), a Perma Pure Inc. Nafion dryer (MD-070), a 2 m long Teflon tube, and then into the  
801 SO<sub>2</sub> analyzer. The initial 1 m of tubing, filter, and drier were located in the humidity-controlled box at  
802 the base of the mast. Dry air was pulled through a charcoal trap and then through the outside of the Nafion



803 dryer at  $2 \text{ L min}^{-1}$ . The analyzer was run with two channels (0 – 20 ppb full scale and 0 – 100 ppb full  
804 scale) and a 20 sec averaging time.

805

806 Zero air (scrubbed with a charcoal trap) was introduced into the sample line upstream of the Fluoropore  
807 filter for 10 min every hour to establish a zero baseline. An  $\text{SO}_2$  standard was generated with a permeation  
808 tube held at  $50^\circ\text{C}$ . The flow over the permeation tube, diluted to 17.7 ppb, was introduced into the sample  
809 line upstream of the Fluoropore filter for 10 min every 6 hrs (Bates et al., 2004).

810

### 811 3.5.3. Radon

812 The rate of emission of radon from the ocean is  $\sim 100$  times less than over land. As a result,  $^{222}\text{Rn}$  is a  
813 qualitative tracer of an air mass that has been recently influenced by continental emissions (Carlson et al.,  
814 1972).

815

816 Radon was measured on all cruises starting with ACE-1 except AEROSOL99, INDOEX, NAURU99,  
817 and DYNAMO. Radon ( $^{222}\text{Rn}$  – half-life of 3.82 days) was measured using the two-filter detector method  
818 of Whittlestone et al. (1998a). Air is drawn through a HEPA filter which removes all radon and thoron  
819 decay products (i.e., daughters), then through a delay chamber in which some daughters are produced.  
820 Finally, the air passes through a second filter which retains the daughters. These daughters have been  
821 produced in controlled conditions so their number is proportional to the radon concentration. A  
822 photomultiplier then counts the radon daughters produced in a 750 L decay tank for a 30-min period. The  
823 detector was standardized using radon emitted from a dry radon source (RN-25, PylonElectronics Corp).  
824 Background counts were measured under conditions of zero air flow (Quinn et al., 2022).

825

### 826 3.5.4. DMS

827 Air and seawater samples for DMS were analyzed using an automated purge and trap system. Air samples  
828 were collected through a Teflon line which ran approximately 60 m from the top of the aerosol sampling  
829 mast to the instrument. One hundred  $\text{mL min}^{-1}$  of the  $4 \text{ L min}^{-1}$  flow were pulled through a KI solution at  
830 the instrument to eliminate oxidant interferences (Cooper et al., 1993). The air sample volume ranged



831 from 0.5 to 1.5 L depending on the DMS concentration. Water vapor was removed by passing the flow  
832 through a  $-25^{\circ}\text{C}$  Teflon tube filled with silanized glass wool. DMS was then trapped in another  $-25^{\circ}\text{C}$   
833 Teflon tube filled with Tenax. During the sample trapping period, methylethyl sulfide (MES) was added  
834 to the sample stream as an internal standard. At the end of the sampling/purge period, the coolant was  
835 pushed away from the trap and the trap was electrically heated. DMS was desorbed onto a DB-1 mega-  
836 bore fused silica column where the sulfur compounds were separated isothermally at  $50^{\circ}\text{C}$  quantified with  
837 either a Flame Photo Detector (FPD) or a Sulfur Chemiluminescence Detector (SCD). The instrument  
838 was calibrated gravimetrically with calibrated permeation tubes. More details of the analysis can be found  
839 in Bates et al. (1998b).

840

### 841 **3.6. Seawater species**

#### 842 **3.6.1. DMS**

843 Seawater samples for DMS analysis were collected from the ship's seawater pumping system at a depth  
844 of approximately 5 m below the ship's waterline. Periodically, a 5 mL water sample was valved from the  
845 ship's water line into a Teflon gas stripper. The sample was purged with hydrogen for 5 min. DMS and  
846 other sulfur gases in the hydrogen purge gas were collected on the Tenax trap held at  $-25^{\circ}\text{C}$  as for the air  
847 samples. Seawater and air sample analysis was identical.

848

#### 849 **3.6.2. $\text{NH}_4^+$ and $\text{NO}_3^-$**

850 Seawater samples for the analysis of  $\text{NH}_4^+$  and  $\text{NO}_3^-$  were taken from a depth of  $\sim 5$  m using the ship's  
851 seawater pumping system. Samples were analyzed for  $\text{NH}_4^+$  using the phenolhypochlorite colorimetric  
852 method of Solarzano (1969) and for  $\text{NO}_3^-$  using the method of Parsons et al. (1984). Both of the analyses  
853 were undertaken with a Technicon Autoanalyzer II (Technicon Corp., Tarrytown, New York).

854

#### 855 **3.6.3. Chlorophyll-a**

856 Discrete seawater samples for chlorophyll-a analysis were taken from the ship's seawater pumping system  
857 2 to 6 times per day. Samples were immediately filtered, put into 10 mL of 90% acetone, and frozen.  
858 Samples were analyzed with a fluorometer within 3 to 4 days onboard the ship. Depending on the cruise,



859 the fluorometer was calibrated several times during, before, or after the experiment usually with algal  
 860 chlorophyll ‘a’ (Sigma Chemical Corp.). The discrete samples were used to calibrate continuous  
 861 fluorescence measurements of seawater also from the ship’s underway seawater pumping system.

862

863 **Table 5. Measurements of gas phase and seawater species on each cruise and the instrumentation used.**

Parameter and Measurement Method	Cruise
O <sub>3</sub> Dasibi 1008-AH UV <sup>a</sup> Photometer	PSI-91, MAGE92, RITS93, RITS94, ACE-1, CSP, ACE-2, AEROSOLS99, INDOEX, ACE-Asia, NEAQS 2004
O <sub>3</sub> TECO Model 49 O <sub>3</sub> Analyzer	RITS94, ACE-1, CSP, ACE-2, AEROSOLS99, INDOEX, ACE-Asia, NEAQS 2002, NEAQS 2004,
O <sub>3</sub> TECO Model 49C O <sub>3</sub> Analyzer	TexAQS, ICEALOT, VOCALS, CalNex, DYNAMO, WACS, WACS-2, NAAMES-1, NAAMES-2, NAAMES-3, NAAMES-4, ATOMIC
SO <sub>2</sub> TEI Model 43C SO <sub>2</sub> Analyzer	ACE-Asia, NEAQS 2002, NEAQS 2004, TexAQS, ICEALOT, VOCALS, CalNex
Radon ( <sup>222</sup> Rn) Two-filter Radon Detector	ACE-1, CSP, ACE-2, ACE-Asia, ACE-Asia, NEAQS 2002, NEAQS 2004, TexAQS, ICEALOT, VOCALS, CalNex, WACS, WACS-2, NAAMES-1, NAAMES-2, NAAMES-3, NAAMES-4, ATOMIC
Atmospheric DMS Purge-and-trap system with FPD <sup>b</sup>	MAGE92, RITS93
Atmospheric DMS Purge-and-trap system with SCD <sup>c</sup>	RITS94, ACE-1, ACE-2, AEROSOLS99, INDOEX, ACE-Asia
Seawater DMS Purge-and-trap system with FPD	PSI-91, MAGE92, RITS93
Seawater DMS Purge-and-trap system with SCD	RITS94, ACE-1, CSP, ACE-2, AEROSOLS99, INDOEX, ACE-Asia, NEAQS 2004, TexAQS, ICEALOT, VOCALS, CalNex, WACS, WACS-2
Seawater NH <sub>4</sub> <sup>+</sup> Technicon Autoanalyzer II	RITS94
Seawater NO <sub>3</sub> - Technicon Autoanalyzer II	PSI-91, RITS93, RITS94, ACE-1, CSP, AEROSOLS99, INDOEX, NAURU99
Seawater Chl-a <sup>d</sup> Fluorometer	PSI-91, RITS93, RITS94, ACE-1, CSP, ACE-2, AEROSOLS99, INDOEX, NAURU99, CalNex, WACS, WACS-2, NAAMES-1, NAAMES-2, NAAMES-3, NAAMES-4, ATOMIC

864 <sup>a</sup>Ultra-Violet

865 <sup>b</sup>Flame Photometric Detector

866 <sup>c</sup>Sulfur Chemiluminescence Detector

867 <sup>d</sup>Chlorophyll-a

868

869

870



### 871 **3.7. Ancillary parameters**

872 Ancillary meteorological and seawater parameters were routinely measured on all cruises. These  
873 parameters include latitude, longitude, ship's speed and course, true wind speed and direction, relative  
874 wind speed and direction, ambient temperature and relative humidity, barometric pressure, and rain rate.  
875 Radiosonde data are available for all cruises except PSI-91, MAGE92, RITS93, RITS94, DYNAMO, and  
876 WACS. Seawater parameters include sea surface temperature and salinity.

877

### 878 **4 Summary of Major Findings**

879 Listed below are many of the high-level major findings reported by PMEL based on its global ocean data  
880 set of marine aerosol properties. Although they may seem fundamental now, they are a result of early  
881 foundational measurements built upon over time with additional cruises in different parts of the world's  
882 oceans. These findings were not developed solely by PMEL but along with other pioneering shipboard  
883 and aircraft measurements made by many other researchers.

884 **1. Measurements of key sulfur species in surface seawater show that most seawater DMS is**  
885 **microbially consumed in the water column, while the ocean-to-atmosphere flux of DMS is a**  
886 **minor sink in the seawater sulfur cycle** (Bates et al., 1994) (Data from PSI-91).

887 **2. The mean surface seawater DMS concentration in the equatorial Pacific (15°S to 15°N) is**  
888 **relatively constant seasonally and interannually.** Large interannual variations associated with  
889 El Nino – Southern Oscillation (ENSO) events appear to have little effect on the concentration of  
890 DMS in tropical surface ocean waters (Bates et al., 1994). (Data from MAGE92, RITS93, RITS94,  
891 ACE-1, CSP and previous cruises not described here).

892 **3. New particle production in the marine boundary layer is rare** due to the high surface area of  
893 sea salt aerosol resulting in the condensation of gas phase precursors onto existing aerosol (Quinn  
894 et al., 1993; Covert et al., 1996; Bates et al., 1998b). (Data from PSI-91, MAGE92, RITS93,  
895 RITS94, ACE-1).

896 **4. The marine aerosol number size distribution has modal characteristics that depend on large**  
897 **scale meteorological features and marine boundary layer residence times.** Strong subsidence  
898 and entrainment from the FT produce an aerosol dominated by particles in the ultra – fine and



- 899 Aitken modes (~ 2 to 80 nm). Residence time in the MBL of a few days or more results in a  
900 bimodal aerosol with Aitken and accumulation modes (80 to 300 nm in diameter) (Covert et al.,  
901 1996; Quinn et al., 1996; Bates et al., 1998b; Bates et al., 2000; Bates et al., 2001; Bates et al.,  
902 2002; Quinn et al., 2017). (Data from MAGE92, RITS93, RITS94, ACE-1, ACE-2, AEROSOL99,  
903 ICEALOT, WACS-2, NAAMES-1).
- 904 **5. Regional and mesoscale meteorological transport patterns impact aerosol number and**  
905 **volume distributions, chemical composition, and optical and cloud-nucleating properties**  
906 (Bates et al., 2001; Bates et al., 2002; Bates et al., 2004; Bates et al., 2008; Quinn et al., 2022).  
907 (Data from ACE-1, ACE-2, AEROSOLS99, ACE-Asia, NEAQS 2002, NEAQS 2004, TexAQS,  
908 ATOMIC).
- 909 **6. Sea salt can comprise a significant mass fraction of not only supermicron but also submicron**  
910 **aerosol in the marine atmosphere. Its relatively large mass concentration, high scattering**  
911 **efficiency, and lifetime comparable to other submicron chemical components often results in**  
912 **submicron sea salt being the dominant contributor to submicron scattering in the marine**  
913 **boundary layer** (Quinn et al., 1996; Quinn et al., 1998b; Murphy et al., 1998; Quinn et al., 1999;  
914 Quinn et al., 2000; Quinn et al., 2005a). (Data from PSI-91, MAGE92, RITS93, RITS94, ACE-1,  
915 ACE-2, AEROSOLS99, INDOEX, ACE-Asia, CSP, NEAQS 2002).
- 916 **7. Instantaneous wind speed often only accounts for a small fraction of the variance in the**  
917 **coarse mode number concentration (~30%) and sea salt submicron and supermicron mass**  
918 **concentrations (~20 to 78%)** due to variability in upwind conditions and advection to the  
919 measurement location (Bates et al., 1998b; Quinn et al., 1999). (Data from PSI-91, MAGE92,  
920 RITS93, RITS94, ACE-1).
- 921 **8. A variable and often large fraction of submicron aerosol mass in the marine boundary layer,**  
922 **both remote and continentally influenced, is not non-sea salt sulfate** (McInnes et al., 1996;  
923 Quinn et al., 2000; Quinn et al., 2005a). (Data from PSI-91, MAGE92, RITS93, RITS94, ACE-1,  
924 ACE-2, AEROSOLS99, INDOEX, ACE-Asia, CSP, NEAQS 2002).
- 925 **9. Sea salt makes up a small fraction of marine boundary layer cloud condensation nuclei**  
926 (Quinn et al., 2017; Quinn et al., 2019). **Instead, the CCN population between 70°S and 80°N**



927 **is composed primarily of nss SO<sub>4</sub><sup>-</sup>** due to large-scale meteorological features that result in  
928 entrainment of particles from the FT into the MBL and regionally varying MBL aerosol residence  
929 times. (Data from RITS93, RITS94, ACE-1, ICEALOT, WACS-2, NAAMES-1, NAAMES-2,  
930 NAAMES-3, NAAMES-4).

931 **10. Particulate organic matter and its degree of oxidation impacts the relative humidity**  
932 **dependence of light scattering and aerosol cloud nucleation** (Quinn et al., 2005b; Quinn et al.,  
933 2008). (Results from INDOEX, ACE-Asia, NEAQS 2002, TexAQS).

934  
935

## 936 **5 Data Set Usage**

937 Examples of previous uses of the data based on PMEL co-authorship are listed below.

938

939 **1. Constraints on models and parameterizations** (e.g., Global distribution of sea salt aerosols:  
940 new constraints from in situ and remote observations (Jaegle et al., 2011); A review of sea-spray  
941 aerosol source functions using a large global set of sea salt aerosol concentration measurements  
942 (Gyrthe et al., 2014); Atmospheric sulfur cycle simulated in the global model GOCART'  
943 Comparison with field observations and regional budgets (Chin et al., 2000); Modelled radiative  
944 forcing of the direct aerosol effect with multi-observation evaluation (Myhre et al., 2009);  
945 Numerical study of Asian dust transport during the springtime of 2001 simulated with the  
946 Chemical Weather Forecasting System (CFORS) model (Uno et al., 2004); CCN predictions using  
947 simplified assumptions of organic aerosol composition and mixing state: a synthesis from six  
948 different locations (Ervens et al., 2010); A model for the radiative forcing during ACE-Asia  
949 derived from CIRPAS Twin Otter and R/V Ronald H. Brown data and comparison with  
950 observations (Conant et al., 2003); Global sea-salt modeling: Results and validation against  
951 multicampaign shipboard measurements (Witek et al., 2007); The Global Aerosol Synthesis and  
952 Science Project (GASSP) (Reddington et al., 2017)).

953



- 954       **2. Intercomparison of instruments and methods** (e.g., ACE-Asia intercomparison of a thermal-  
955       optical method for the determination of particle-phase organic and elemental carbon (Schauer et  
956       al., 2003); Bias in filter based aerosol absorption measurements due to organic aerosol loading:  
957       Evidence from ambient measurements (Lack et al., 2008); Comparison of aerosol single scattering  
958       albedos derived by diverse techniques in two North Atlantic experiments (Russell et al., 2002)).  
959
- 960       **3. Comparison to and validation of remote retrievals** (e.g., Measurements of aerosol vertical  
961       profiles and optical properties during INDOEX 1999 using micropulse lidars (Welton et al., 2002);  
962       Geostationary satellite retrievals of aerosol optical thickness during ACE-Asia (Wang et al.,  
963       2003); Clear-sky infrared aerosol radiative forcing at the surface and the top of the atmosphere  
964       (Markovic et al., 2003); Spectral absorption of solar radiation by aerosols during ACE-Asia  
965       (Bergstrom et al., 2004); Lidar measurements during Aerosols99 (Voss et al., 2001); Multi-grid-  
966       cell validation of satellite aerosol property retrievals in INTEX/ITCT/ICARTT 2004 (Russell et  
967       al., 2007); Shipboard sunphotometer measurements of aerosol optical depth spectra and columnar  
968       water vapor during ACE 2 and comparison with selected land, ship, aircraft, and satellite  
969       measurements ).  
970
- 971       **4. Addition to larger data sets** (e.g., Maritime aerosol network as a component of aerosol robotic  
972       network (Smirnov et al., 2009); Total observed organic carbon (TOOC) in the atmosphere: a  
973       synthesis of North American observations (Heald et al., 2008); A global database of sea surface  
974       DMS measurements and a simple model to predict sea surface DMS as a function of latitude,  
975       longitude and month (Kettle et al., 1999)).  
976

## 977 **6 Summary**

978 PMEL conducted 25 cruises between 1991 and 2020 measuring aerosol chemical, microphysical, optical,  
979 and cloud nucleating properties. These cruises provide coverage in all of the world's oceans resulting in  
980 the largest global ocean data set of marine aerosol properties. The data set also includes gas phase and  
981 seawater species. PMEL's major findings and data usage by others are summarized. A description of each



982 cruise is provided including location, timing, and objectives. References are cited to provide a deeper  
 983 context for each cruise. The intention of the paper is to advance widespread use of the data by the broader  
 984 research community.

985

## 986 **7 Data availability**

987 All cruise data sets are publicly available at NOAA’s National Centers for Environmental Information  
 988 (<https://www.ncei.noaa.gov>). Cruise identification, data links (DOIs), and data references are provided  
 989 in Table 6. The data are permanently and publicly available at NCEI.

990

991 ***Table 6. Summary of cruise data links (DOIs), and references. See Table 1 for dates, ports, and ocean***  
 992 ***region for each cruise. The data are permanently and publicly available at NOAA’s NCEI.***

993

Cruise	Data links	Data reference
PSI-91	<a href="https://doi.org/10.25921/44nn-d608">https://doi.org/10.25921/44nn-d608</a>	(Quinn et al., 2026i)
MAGE92	<a href="https://doi.org/10.25921/bz8f-b917">https://doi.org/10.25921/bz8f-b917</a>	(Quinn et al., 2026g)
RITS93	<a href="https://doi.org/10.25921/ec4p-9410">https://doi.org/10.25921/ec4p-9410</a>	(Quinn et al., 2026f)
RITS94	<a href="https://doi.org/10.25921/ec4p-9410">https://doi.org/10.25921/ec4p-9410</a>	(Quinn et al., 2026f)
ACE-1 Leg 1	<a href="https://doi.org/10.25921/z3bm-y330">https://doi.org/10.25921/z3bm-y330</a>	(Quinn et al., 2026c)
ACE-1 Leg 2	<a href="https://doi.org/10.25921/z3bm-y330">https://doi.org/10.25921/z3bm-y330</a>	(Quinn et al., 2026c)
CSP	<a href="https://doi.org/10.25921/pgzy-5h08">https://doi.org/10.25921/pgzy-5h08</a>	(Quinn et al., 2026b)
ACE-2	<a href="https://doi.org/10.25921/3fk0-0m36">https://doi.org/10.25921/3fk0-0m36</a>	(Quinn et al., 2025e)
AEROSOLS99	<a href="https://doi.org/10.25921/67kx-2d82">https://doi.org/10.25921/67kx-2d82</a>	(Quinn et al., 2026h)
INDOEX	<a href="https://doi.org/10.25921/67kx-2d82">https://doi.org/10.25921/67kx-2d82</a>	(Quinn et al., 2026h)
NAURU99	<a href="https://doi.org/10.25921/e2rz-yg88">https://doi.org/10.25921/e2rz-yg88</a>	(Quinn et al., 2026k)
ACE-Asia	<a href="https://doi.org/10.25921/jd13-t245">https://doi.org/10.25921/jd13-t245</a>	(Quinn et al., 2026l)
NEAQS 2002	<a href="https://doi.org/10.25921/q66h-r438">https://doi.org/10.25921/q66h-r438</a>	(Quinn et al., 2026j)
NEAQS 2004	<a href="https://doi.org/10.25921/q66h-r438">https://doi.org/10.25921/q66h-r438</a>	(Quinn et al., 2026j)
TexAQS-GoMACCS	<a href="https://doi.org/10.25921/c6n1-0840">https://doi.org/10.25921/c6n1-0840</a>	(Quinn et al., 2025d)
ICEALOT	<a href="https://doi.org/10.25921/bgy4-3075">https://doi.org/10.25921/bgy4-3075</a>	(Quinn et al., 2025c)



VOCALS	<a href="https://doi.org/10.25921/mafz-2n04">https://doi.org/10.25921/mafz-2n04</a>	(Quinn et al., 2025f)
CalNex	<a href="https://doi.org/10.25921/xf4m-dx08">https://doi.org/10.25921/xf4m-dx08</a>	(Quinn et al., 2025a)
DYNAMO	<a href="https://doi.org/10.25921/m0ec-rn58">https://doi.org/10.25921/m0ec-rn58</a>	(Quinn et al., 2026e)
WACS	<a href="https://doi.org/10.25921/tx5t-1e17">https://doi.org/10.25921/tx5t-1e17</a>	(Quinn et al., 2026d)
WACS2	<a href="https://doi.org/10.25921/tx5t-1e17">https://doi.org/10.25921/tx5t-1e17</a>	(Quinn et al., 2026d)
NAAMES1	<a href="https://doi.org/10.25921/df6d-p183">https://doi.org/10.25921/df6d-p183</a>	(Quinn et al., 2025b)
NAAMES2	<a href="https://doi.org/10.25921/df6d-p183">https://doi.org/10.25921/df6d-p183</a>	(Quinn et al., 2025b)
NAAMES3	<a href="https://doi.org/10.25921/df6d-p183">https://doi.org/10.25921/df6d-p183</a>	(Quinn et al., 2025b)
NAAMES4	<a href="https://doi.org/10.25921/df6d-p183">https://doi.org/10.25921/df6d-p183</a>	(Quinn et al., 2025b)
ATOMIC	<a href="https://doi.org/10.25921/w7ab-3s87">https://doi.org/10.25921/w7ab-3s87</a>	(Quinn et al., 2026a)

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997 **Author contributions**

998 T.S.B. and P.K.Q. conceptualized research goals. P.K.Q., T.S.B., D.J.C., J.E.J., and L.M.U. participated  
999 in collecting and analyzing data. P.K.Q. prepared the paper. D.J.C. and H.B. prepared data sets for  
1000 archival at NCEI.

1001

1002 **Competing interests**

1003 The authors declare that they have no conflict of interest.

1004

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1017 **References**

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