



1 GSSM-10 (Global 10-m Surface Soil Moisture) Derived from 2 Multi-Sensor Data and Ensemble Learning

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9

10 Abstract

11 Satellite-driven soil moisture monitoring systems currently available fail to meet the
 12 spatial resolution requirement for a wide range of applications. This limitation is
 13 particularly critical for agricultural water management, assessing risks associated with
 14 extreme events, and hydrological modeling. This work aims to address the spatial
 15 limitations of satellite soil moisture remote sensing by developing GSSM-10, a global
 16 10-meter resolution surface soil moisture dataset, using multi-sensor datasets
 17 integrated within an ensemble machine learning framework. These datasets
 18 encompass diverse data types—active microwave, multispectral, thermal infrared, and
 19 land elevation—offering a robust and comprehensive approach to estimating surface
 20 soil moisture (SSM). The ensemble model incorporates TabNet, Random Forest (RF),
 21 and Extreme Gradient Boosting (XGBoost). The model was trained on ground-truth
 22 data collected from the International Soil Moisture Network (ISMN). The ensemble
 23 model demonstrated robust performance, achieving an R^2 of 0.8344, a bias of -0.0001 ,
 24 an RMSE of $0.0433 \text{ m}^3/\text{m}^3$, and an ubRMSE of $0.0433 \text{ m}^3/\text{m}^3$ in 5-fold
 25 cross-validation. When evaluated on a held-out test set, the model achieved similar
 26 levels of accuracy, with an R^2 of 0.8591, a bias of $-0.0002 \text{ m}^3/\text{m}^3$, and an
 27 RMSE/ubRMSE of $0.0401 \text{ m}^3/\text{m}^3$. An interactive web platform has been developed
 28 for data access, visualization, and download, enabling broad adoption by researchers,
 29 practitioners, and policymakers. By providing globally consistent, high-resolution SM
 30 estimates, GSSM-10 represents a significant advancement in satellite-based soil
 31 moisture monitoring for environmental and agricultural applications.

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33 **Keywords:** soil moisture, remote sensing, data fusion, machine learning.

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37 1. Introduction



1 Soil moisture plays a pivotal role in the Earth's water and energy cycles, influencing
2 climate, hydrology, and ecosystem dynamics (Klein and Taylor, 2020; Porkka et al.,
3 2024; Sun et al., 2025; Zhang et al., 2020). It governs the partitioning of rainfall into
4 runoff or infiltration and modulates evapotranspiration, thereby affecting weather and
5 climate patterns at regional to global scales (Denissen et al., 2022; Seneviratne et al.,
6 2010; Sun et al., 2025). Recognized as an essential climate variable, soil moisture data
7 are indispensable for climate research (Humphrey et al., 2021; Liu et al., 2023; Qiao
8 et al., 2023; Seneviratne et al., 2010; Soares and Lima, 2022), hydrological modeling
9 (Droppers et al., 2024; Eini et al., 2023; Fatima et al., 2024; Leonarduzzi et al., 2021;
10 Mei et al., 2023), drought and flood forecasting (Lesinger and Tian, 2025; Qing et al.,
11 2023; Wasko and Nathan, 2019; Wyatt et al., 2020; Yao et al., 2023), and agricultural
12 management (Chatterjee et al., 2022; Li et al., 2022; Martínez-Fernández et al., 2016;
13 McNairn et al., 2012; Zhou et al., 2021).

14
15 Compared to traditional in-situ point measurement soil moisture (SM) sensors,
16 satellite remote sensing offers cost-effective and large-scale SM monitoring solutions
17 (Chaudhary et al., 2022; Cheng et al., 2022; Dubois et al., 2021). NASA's Soil
18 Moisture Active Passive (SMAP) (Entekhabi et al., 2010), EUMETSAT's Advanced
19 Scatterometer (ASCAT) (Wagner et al., 2013), and ESA's Soil Moisture and Ocean
20 Salinity (SMOS) (Kerr et al., 2001) have demonstrated the power of spaceborne
21 sensors to map surface soil moisture globally. However, their low spatial resolutions
22 are not pertinent to many applications, particularly to agriculture and field-level scale
23 applications (Babaeian et al., 2021; Nguyen et al., 2022). For example, SMAP's
24 radiometer measures moisture at ~36 km resolution (with an enhanced 9 km product),
25 ASCAT scatterometer products are available at ~12.5 km, and SMOS provides data at
26 ~25 km. Even multi-sensor merged datasets like the ESA Climate Change Initiative
27 (CCI) soil moisture (which merges multiple satellites) are typically gridded at 0.25°
28 (~25 km) (Dorigo et al., 2017). At these scales, crucial fine-scale heterogeneity is lost.
29 Model-based products such as ERA5-Land (Muñoz-Sabater et al., 2021), which
30 provides hourly global soil moisture data at ~9 km resolution, and GLDAS (Global
31 Land Data Assimilation System), which produces 3-hourly estimates at 0.25°
32 resolution (Syed et al., 2008; Zawadzki and Kędzior, 2016), offer valuable temporal
33 continuity but remain too coarse for capturing sub-field heterogeneity. While these
34 datasets are essential for large-scale hydrological modeling and climate analysis, their
35 spatial granularity is insufficient for localized applications (Liu et al., 2019; Xu et al.,
36 2021). This mismatch between the scale of observation and the scale of
37 decision-making severely constrains the usability of current global datasets for
38 applications such as irrigated agriculture, flood forecasting, and wildfire risk
39 assessment (Gebrechorkos et al., 2023; He et al., 2023; Peng et al., 2021; Sabaghy et
40 al., 2020).

41
42 A variety of global SM datasets have been developed using remote sensing, land
43 surface models, and machine learning. Recently, several products with a resolution of
44 ~1 km have emerged, marking a significant step toward finer spatial detail. For



1 example, Fan et al. (2025) introduced a global 1-km SM product derived from
2 Sentinel-1 SAR observations. Zhang et al. (2023) generated a daily 1-km global
3 surface SM dataset for 2000–2020 by integrating multi-source satellite-driven
4 information (albedo, land temperature, leaf area index) and reanalysis data using an
5 ensemble learning (XGBoost) model. Zheng et al., 2023 achieved a similar 1-km
6 global product by downscaling the 0.25° ESA CCI satellite soil moisture with
7 Random Forest, producing a gap-free daily 21-year record (2000–2020). Han et al.,
8 2023 developed another 1-km global dataset (surface top 5 cm layer) using a
9 physics-informed machine learning approach, achieving high accuracy (correlation
10 coefficient of ~0.9) from 2000 to 2020. These high-resolution global datasets
11 represent significant progress in capturing soil moisture at much finer scales than
12 earlier global products. Despite their improved resolution, the ~1 km global datasets
13 still fall short for applications demanding sub-field-scale detail. A 1-km pixel
14 (~100 ha) averages over heterogeneous terrain and management units and is too
15 coarse for field-level precision applications.

16
17 To meet fine-scale needs, several regional datasets push spatial resolutions to tens or
18 hundreds of meters, albeit over limited areas. Vergopalan et al. (2021) developed
19 SMAP-HydroBlocks, a 30-m resolution surface 5cm soil moisture dataset for the
20 conterminous U.S. (2015–2019) by integrating high-resolution land surface modeling
21 with downscaled satellite observations and machine learning. Baghdadi et al., (2017)
22 developed the S2MP (Sentinel-1/Sentinel-2-derived Soil Moisture Product), a regional
23 10 m resolution soil moisture dataset. However, its coverage is presently limited to
24 specific regions (e.g., parts of southern France). In summary, high spatial resolution
25 satellite-derived soil moisture is available at regional scales, but current datasets are
26 constrained by both limited geographic coverage and relatively short temporal records
27 (Table 1). Therefore, there is a need to develop a soil moisture product that combines
28 global coverage with high spatial resolution.

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Table 1 Information of common soil moisture products.

Category	Datasets	Scale	Period	Spatial Resolution	Temporal Resolution	Depth (cm)
Satellite data	SMAP	Global	2015 - present	9 km	2 - 3 days	0-5; 0-100
	ASCAT	Global	2000 - present	12.5 / 25 km	Twice per day	0-2
	SMOS	Global	2000 - present	35 - 50 km	2 - 3 days	0-5
Artificially developed products	ERA5	Global	1950 - present	11.132 Km	hourly / daily	0-7, 7-28, 28-100, and 100-200
	GLDAS	Global	1948 - present	27.830 Km	daily	0-10,; 10-40, 40-100, and



						100-200
Skulovich and Gentine, 2023	Global	2002 - 2020	25 km	3days	0-5	
Han et al., 2023	Global	2000 - 2020	1 km	daily	0-5	
Zheng et al., 2023	Global	2000 - 2020	1 km	daily	0-5	
Zhang et al., 2023	Global	2000 - 2020	1 km	daily	0-5	
O. and Orth, 2021	Global	2000 - 2019	0.25° (27.75 km)		0-10, 10-30, and 30-50	
SMRFR (Liu et al., 2025)	Global	2000 - 2023	9 km	daily	0-5, 5-10, 10-30, 30-50, and 50-100	
SMAP-HydroBlocks (Vergopolan et al., 2021)	U.S.	2015 - 2019	30 m	6 hours	0-5	
Cui et al., 2019	Tibet Plateau	2002 - 2015	0.25° (27.75 km)	daily	0-5	
Song et al., 2022	China	2003 - 2019	1 km	daily	0-10	
S2MP (Lozac'h et al., 2020)	parts of France, Morocco, Germany, U. S. A	2017 - 2024	field scale	Synchronize with Sentinel-1/2	0-5	

To address existing limitations in soil moisture monitoring, this research introduces a global 10-m surface soil moisture dataset developed using multi-sensor remote sensing data, including active microwave, multispectral, thermal, and elevation inputs, and advanced ensemble machine learning techniques. The dataset provides continuous near real-time updates and includes historical records dating back to January 2016. We employed a variety of ensemble models, such as TabNet (Arik and Pfister, 2021), Random Forest (Belgiu and Drăguț, 2016), and XGBoost (Chen and Guestrin, 2016), to reduce individual model bias and leverage their complementary strengths. This work represents a significant advancement in high-resolution soil moisture mapping, with broad applications in precision agriculture, high-resolution disaster risk assessment, sub-catchment hydrological modeling, and (micro-) climate research.

2. Methodology

2.1 Data



2.1.1 Remote Sensing Data

In this study, we utilized a variety of datasets, each with a unique role in soil moisture analysis (Table 2). Sentinel-1 with Synthetic Aperture Radar is capable of emitting microwaves that penetrate the surface, with varying degrees of signal reflection indicating different levels of moisture in the soil, even under vegetation cover.

Sentinel-2 and Landsat-8/9 provide multispectral imagery highlighting information about vegetation and moisture using various vegetation indices such as the Normalized Difference Vegetation Index (NDVI, eq. (1)) (Pettorelli et al., 2005; Xu et al., 2025) and the Normalized Difference Moisture Index (NDMI, eq. (2)) (Mkhwenkwana et al., 2025; Monteiro et al., 2024). Both datasets were used jointly to enrich the spectral feature space. Differences in sensor characteristics, spectral band configurations, and observation geometries provide complementary information that enhances the model's ability to generalize across varying land surface and atmospheric conditions. Moreover, land topography obtained from ALOS DSM and STRM DEM has a significant impact on soil moisture distribution. For example, areas at higher elevations may experience more runoff, resulting in drier soils, while areas at lower elevations may retain more moisture. ALOS DSM is the primary elevation data used. In areas where ALOS DSM is unavailable, the SRTM DEM is used as a substitute.

$$NDVI = \frac{(NIR - Red)}{(NIR + Red)} \quad (1)$$

$$NDMI = \frac{(NIR - SWIR)}{(NIR + SWIR)} \quad (2)$$

Table 2. Summary of the datasets used in the study.

Sensor	Spatial resolution	Temporal resolution	Features	Time
Sentinel-1	10 m	3-5 days	VV, VH, incident angle	1/1/2016-12/31/2023
Sentinel-2A/B	10 m	2-3 days	B1 - B12, NDVI, NDMI	1/1/2016-12/31/2023
Landsat-8/9	30	8 days	B1 - B11, NDVI, NDMI	1/1/2016-12/31/2023
ALOS DSM / SRTM DEM	30 m	Static	Elevation	2006/2000
ISMN	Point-based (station)	Daily	Soil moisture, latitude, longitude	1/1/2016-12/31/2023

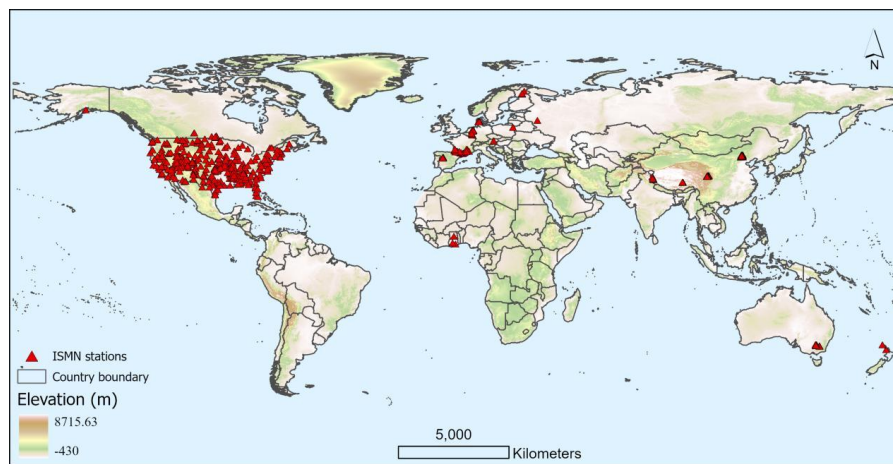
2.1.2 Ground-based soil moisture dataset

We used ground data from the International Soil Moisture Network (ISMN, <https://ismn.earth/en/>) for model training and validation. ISMN is a global in-situ data hosting service that consolidates soil moisture measurements from various sources and networks. We filtered the stations to include only those with available data from



1 January 1, 2016, to December 31, 2023, and provided surface soil moisture
 2 measurements for the top 5 cm of soil depth. In total, data from 699 stations were
 3 used in this study. The spatial distribution of the selected ISMN stations used in this
 4 study is shown in Figure 1. While a large proportion of the selected stations are
 5 located within the United States, the selected stations span diverse climatic conditions
 6 and land cover types.

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9 Fig. 1 Spatial distribution of the selected ISMN stations used in this study.

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12 2.2.1 Water cloud model

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14 To correct for vegetation effects on the radar backscatter prior to soil moisture
 15 retrieval, a dynamic Water Cloud Model (WCM) approach was applied (Attema and
 16 Ulaby, 1978). The WCM relates the observed radar backscatter (σ_{total}) to
 17 contributions from vegetation and soil (Li et al., 2024; Nijaguna et al., 2023).

18

19 For each Sentinel-1 acquisition date, the closest available cloud-free Sentinel-2 image
 20 within a ± 3 -day window was identified to capture the most representative vegetation
 21 conditions at the time of radar observation. From the selected Sentinel-2 image, the
 22 NDVI was computed and sampled at the station location. The sampled NDVI value
 23 was then used to dynamically calculate the WCM parameters, specifically the
 24 vegetation scattering coefficient (A) and attenuation coefficient (B), for both VV and
 25 VH polarizations. Using these dynamically adjusted parameters, vegetation effects
 26 were removed from the Sentinel-1 VV and VH backscatter according to the standard
 27 WCM formulation (Attema and Ulaby, 1978) Finally, the corrected soil surface
 28 backscatter (σ_{soil}) for both polarizations was retained for subsequent analysis and soil
 29 moisture retrieval.

30

$$\sigma_{total} = \sigma_{veg} + \tau^2 \times \sigma_{soil} \quad (3)$$



1

2 where σ_{veg} is the direct vegetation scattering contribution, τ^2 is the two-way
 3 attenuation factor through the vegetation layer, and σ_{soil} is the soil surface
 4 backscatter.

5

6 The two-way attenuation τ^2 was modeled as:

$$7 \quad \tau^2 = \exp(-2B \times V \times \sec\theta) \quad (4)$$

8

9 and the vegetation scattering σ_{veg} was modeled as:

$$10 \quad \sigma_{veg} = A \times V \times \cos\theta \times (1 - \tau^2) \quad (5)$$

11

12 Where V is the vegetation descriptor, here taken as NDVI, θ is the local incidence
 13 angle (in radians), and A and B are empirical coefficients dependent on vegetation
 14 density.

15

16 The soil backscatter was then retrieved by rearranging the WCM equation:

$$17 \quad \sigma_{soil} = \frac{\sigma_{total} - \sigma_{veg}}{\tau^2} \quad (6)$$

18

19 All backscatter computations were performed on a linear scale (i.e., reflectivity
 20 coefficient, unitless), with Sentinel-1 backscatter converted from decibels (dB) to a
 21 linear scale prior to WCM correction and then converted back to dB afterward.

22 Unlike conventional methods that use static WCM parameters, this study dynamically
 23 adjusted the WCM scattering (A) and attenuation (B) coefficients based on NDVI at
 24 each Sentinel-1 acquisition time (Baghdadi et al., 2019, 2017; Rawat et al., 2021).
 25 This allowed the vegetation contribution to vary spatially and temporally in response
 26 to actual plant growth stages.

27

28 For VV polarization:

$$29 \quad A_{VV} = \begin{cases} 0.12 \times NDVI & \text{if } NDVI \leq 0.8 \\ 0.095 \times NDVI & \text{if } NDVI > 0.8 \end{cases} \quad (7)$$

30

31

$$32 \quad B_{VV} = \begin{cases} 0.70 \times NDVI & \text{if } NDVI \leq 0.8 \\ 0.56 \times NDVI & \text{if } NDVI > 0.8 \end{cases} \quad (8)$$

33 For VH polarization:

$$34 \quad A_{VH} = \begin{cases} 0.05 \times NDVI & \text{if } NDVI \leq 0.8 \\ 0.04 \times NDVI & \text{if } NDVI > 0.8 \end{cases} \quad (9)$$

35



$$B_{VH} = \begin{cases} 1.45 \times NDVI & \text{if } NDVI \leq 0.8 \\ 1.16 \times NDVI & \text{if } NDVI > 0.8 \end{cases} \quad (10)$$

Distinct WCM parameters were applied for NDVI values above and below the 0.8 threshold. These dynamic parameterizations allowed the vegetation scattering and attenuation effects to vary smoothly with vegetation conditions.

2.2.2 Ensemble learning

The ISMN and remote sensing data were compiled into a structured tabular format, where each row represents a specific location and acquisition time, and each column corresponds to a predictor variable or the observed soil moisture value. To effectively capture the complex, nonlinear relationships between multi-source remote sensing features and in-situ soil moisture observations, this study employed three complementary machine learning models: TabNet, Random Forest, and XGBoost. TabNet is a deep learning architecture specifically designed for tabular datasets (Arik and Pfister, 2021). It employs attention-based feature selection and sparse representation to enhance interpretability and effectively model complex patterns in high-dimensional tabular data derived from remote sensing (Khaliq et al., 2025; Triana-Martinez et al., 2025). Random Forest (Breiman, 2001), an ensemble of decision trees, offers high robustness to overfitting, effective modeling of nonlinear relationships, and strong performance in the presence of noise and collinearity (Belgiu and Drăguț, 2016; Liu et al., 2025). XGBoost (Chen and Guestrin, 2016), a regularized gradient boosting framework, has demonstrated superior predictive accuracy across a wide range of applications. Its advantages include efficient handling of missing data, built-in regularization, and the capacity to capture complex feature interactions through additive model training (Aydin and Ozturk, n.d.; Deng and Lumley, 2024; Karthikeyan and Mishra, 2021).

The SM estimation workflow is illustrated in Figure 2. We trained TabNet, Random Forest, and XGBoost models using remote sensing features and ISMN ground observations, with 80% of the entire dataset randomly allocated for training and 20% for testing. A 5-fold cross-validation was performed on the entire training dataset.

To optimize the hyperparameters of each model, we employed Optuna, an open-source hyperparameter optimization framework that uses a flexible and efficient sampling-based approach to automate tuning (Akiba et al., 2019). In addition to individual model optimization, Optuna was used to determine the optimal ensemble weights for combining the outputs of TabNet, XGBoost, and Random Forest, thereby integrating TabNet's deep learning feature extraction, XGBoost's gradient boosting, and Random Forest's handling of nonlinear relationships.

Model accuracy was assessed using the coefficient of correlation (R^2 , eq(11)), bias (eq(12)), root mean square error (RMSE, eq(13)), and unbiased root mean square error (ubRMSE, eq(14)). Finally, the optimized ensemble model was used to generate



1 10-m resolution SM maps from January 2016 onward, with real-time updates.

2

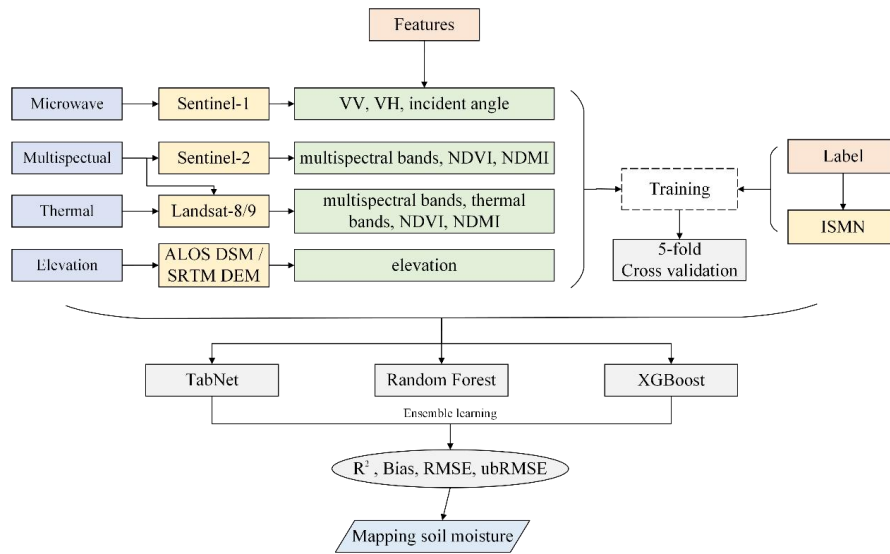
$$3 \quad R^2 = 1 - \frac{\sum (Y_i - \hat{Y}_i)^2}{\sum (Y_i - \bar{Y})^2} \quad (11)$$

$$4 \quad Bias = \frac{1}{n} \sum_{i=1}^n (\hat{Y}_i - Y_i) \quad (12)$$

$$5 \quad RMSE = \sqrt{\frac{1}{n} \sum_{i=1}^n (\hat{Y}_i - Y_i)^2} \quad (13)$$

$$6 \quad ubRMSE = \sqrt{RMSE^2 - Bias^2} \quad (14)$$

7 Where, $\hat{Y}_i$ is the predicted value, Y_i is the observed value, $\bar{Y}$ is the mean of predicted
 8 values, $\bar{Y}$ is the mean of observed values, n is the number of observations.



9

10 Fig. 2. Workflow of SM estimation from multi-source datasets.

11

12 2. Results

13

14 3.1 Model performance

15

16 Table 3 presents the average accuracy metrics obtained from 5-fold cross-validation
 17 for each model. Among the individual models, TabNet achieved the highest
 18 performance, with an R^2 of 0.7756 and the lowest RMSE ($0.0506 \text{ m}^3/\text{m}^3$), followed
 19 closely by Random Forest ($R^2 = 0.7702$, $RMSE = 0.0510 \text{ m}^3/\text{m}^3$). XGBoost showed
 20 comparatively lower accuracy, with an R^2 of 0.6792 and a higher RMSE of 0.0603
 21 m^3/m^3 . In contrast, the ensemble model outperformed all individual models, achieving
 22 the highest R^2 (0.8344) and the lowest RMSE ($0.0433 \text{ m}^3/\text{m}^3$), as well as near-zero
 23 bias and the lowest unbiased RMSE (ubRMSE). These results highlight the advantage

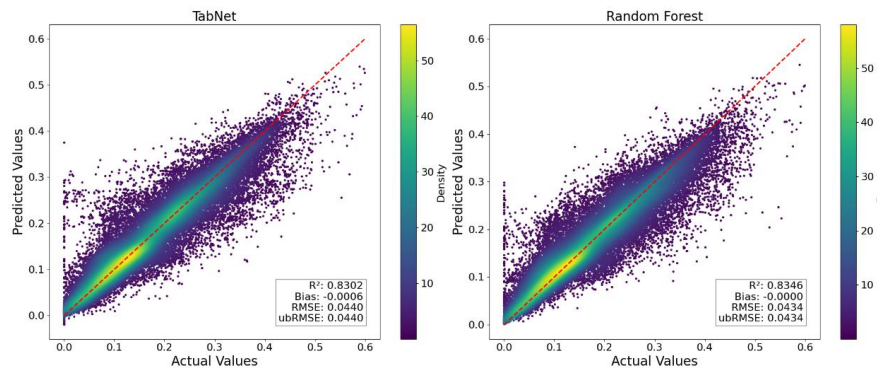


of ensemble learning in integrating the complementary strengths of deep learning (TabNet), decision trees (Random Forest), and gradient boosting (XGBoost) to enhance the accuracy and robustness of soil moisture predictions.

Table 3. Average accuracy metrics from 5-fold cross-validation for each model.

Model	R ²	Bias	RMSE (m ³ /m ³)	ubRMSE (m ³ /m ³)
TabNet	0.7756	-0.0003	0.0506	0.0506
RF	0.7702	0.0002	0.0510	0.0510
XGBoost	0.6792	0.0001	0.0603	0.0603
Ensemble	0.8344	-0.0001	0.0433	0.0433

Predicted versus actual soil moisture values on the 20% test set demonstrate strong agreement across all models, with the ensemble model showing the highest accuracy (Figure 3). The ensemble predictions were generated using a weighted combination of the three models, with weights of 0.56 for Random Forest, 0.43 for TabNet, and 0.26 for XGBoost. The ensemble model achieved an R² of 0.8591, a near-zero bias (−0.0002 m³/m³), and the lowest RMSE and ubRMSE (both 0.0401 m³/m³), indicating excellent predictive performance with minimal systematic error. XGBoost followed closely with an R² of 0.8586 m³/m³ and RMSE of 0.0401 m³/m³, while Random Forest and TabNet yielded slightly lower accuracies (R² = 0.8346 m³/m³ and 0.8302 m³/m³; RMSE = 0.0434 m³/m³ and 0.0440 m³/m³, respectively). However, the ensemble model exhibited the tightest clustering and lowest dispersion, suggesting that it effectively leveraged the complementary predictive capabilities of its constituent algorithms.



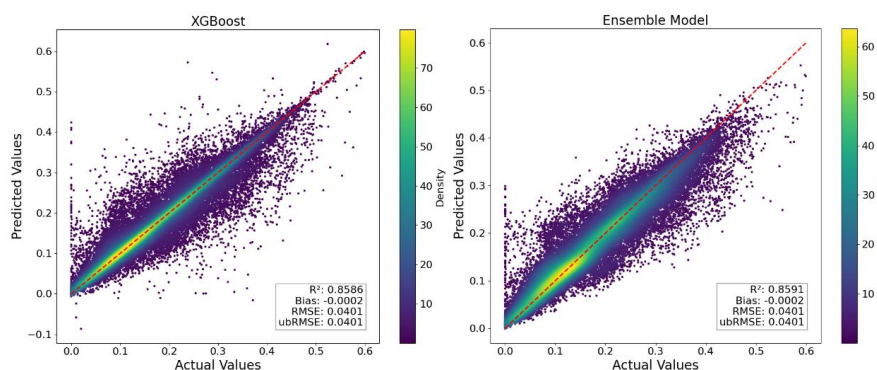


Fig. 3 Scatter density plots of predicted versus actual surface soil moisture on the test set for (a) TabNet, (b) Random Forest, (c) XGBoost, and (d) the Ensemble model.

3.2 Feature importance

Feature importance analysis was conducted across the three individual models—Random Forest, TabNet, and XGBoost—as well as their ensemble to better understand the contributions of each input variable to soil moisture prediction (Figure 4). Across all models, the Digital Surface Model (DSM) emerged as the most influential predictor, highlighting the dominant role of topography in governing soil moisture dynamics. Longitude and latitude were also consistently ranked among the top features, particularly in the ensemble and Random Forest models. These spatial coordinates likely serve as proxies for geospatial trends associated with climate zones, soil texture, and land management practices.

The ensemble model combined predictions from TabNet, Random Forest, and XGBoost using weighted contributions. It emphasized DSM and geolocation (longitude and latitude) as the three most important features, suggesting strong influence from landscape-driven and region-specific moisture patterns. Following these, vegetation-related indicators such as Landsat-based NDMI and NDVI were ranked highly, capturing the relationship between plant water content and soil moisture. A range of spectral reflectance bands from both Landsat (e.g., B3, B4, B11) and Sentinel-2 (e.g., B4, B8, B11) were also assigned substantial importance, reflecting their utility in detecting surface wetness and vegetation vigor. While SAR-derived features (e.g., VV and VH backscatter) were generally ranked lower in the ensemble, their inclusion still contributed valuable information in certain contexts, especially under low vegetation cover.

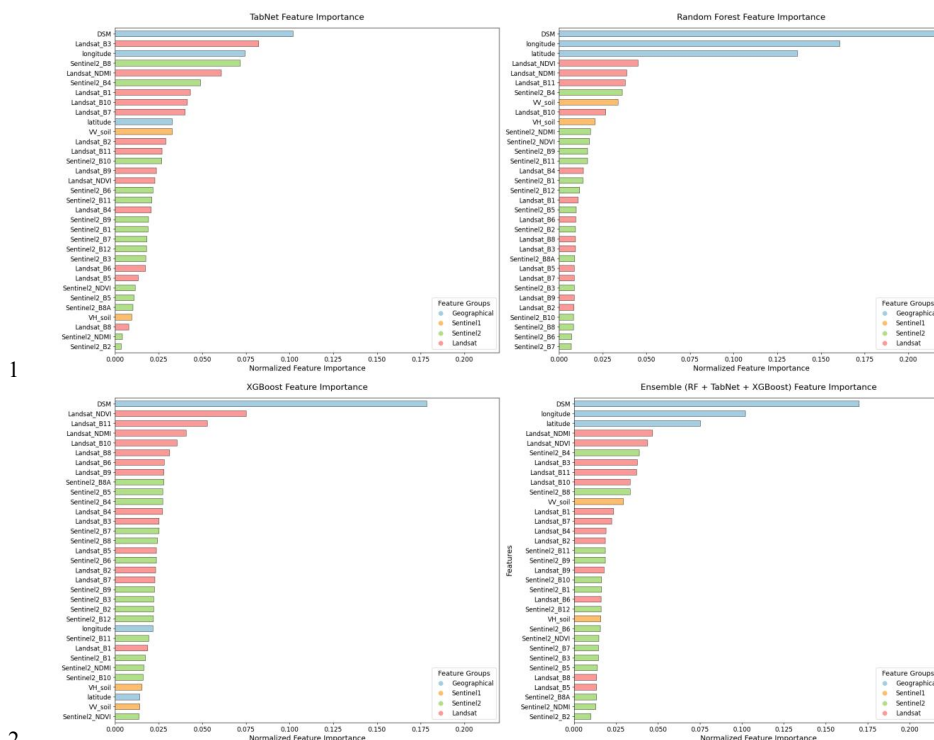


Fig. 4 Feature importance of each model.

3.3 Comparison to other soil moisture datasets

3.3.1 Compare to SMAP-HydroBlocks

The SMAP-HydroBlocks (SMAP-HB) dataset is a high-resolution, satellite-based surface soil moisture product developed at 30 m spatial resolution for the conterminous United States from 2015 to 2019. It was generated by integrating high-resolution land surface modeling, radiative transfer modeling, machine learning, and SMAP satellite microwave observations through a scalable cluster-based merging framework. The dataset was validated using measurements from 1,192 in situ observational sites. SMAP-HB achieved a median temporal correlation of 0.73 ± 0.13 and a median Kling-Gupta Efficiency (KGE) of 0.52 ± 0.20 , indicating marked improvements relative to the standard SMAP Level 3 products.

To assess the relative accuracy of the proposed GSSM-10 dataset, both GSSM-10 and SMAP-HB were evaluated against independent in situ soil moisture measurements obtained from the ISMN. The evaluation metrics summarized in Table 5. The GSSM-10 dataset demonstrated strong agreement with ground observations, achieving an R^2 of 0.8601, a near-zero bias of $-0.0003 \text{ m}^3/\text{m}^3$, and RMSE and ubRMSE values of $0.0406 \text{ m}^3/\text{m}^3$. In contrast, the SMAP-HB dataset exhibited poor



correspondence with the in-situ measurements, yielding a negative R^2 of -0.4253 , a bias of $-0.0075 \text{ m}^3/\text{m}^3$, and RMSE and ubRMSE values of $0.1296 \text{ m}^3/\text{m}^3$ and $0.1293 \text{ m}^3/\text{m}^3$, respectively.

These findings suggest that GSSM-10 offers substantially improved predictive performance over SMAP-HB when benchmarked against independent field observations. The low bias and minimal random error observed in GSSM-10 indicate its robustness in capturing surface soil moisture variability.

Table 4. Validation of SMAP-HB and GSSM-10 using 2569 in situ soil moisture observations from ISMN.

Location	R^2	Bias (m^3/m^3)	RMSE (m^3/m^3)	ubRMSE (m^3/m^3)
GSSM-10	0.8601	-0.0003	0.0406	0.0406
SMAP-HB	-0.4253	-0.0075	0.1296	0.1293

3.3.1 Compared to $S^2\text{MP}$

The $S^2\text{MP}$ (Sentinel-1/Sentinel-2-derived soil moisture product), developed by Baghdadi et al. (2017), was designed to estimate surface soil moisture at the plot scale by coupling radar backscatter from Sentinel-1 with vegetation indices derived from Sentinel-2 imagery using a neural network inversion approach. The product has been validated against in situ measurements collected across several regions, including parts of France, Morocco, Germany, and the United States. It achieved a RMSE of approximately 5 vol.%, demonstrating a high level of accuracy in agricultural regions with vegetation cover. In addition to ground-based validation, $S^2\text{MP}$ has been compared with other widely used soil moisture products, including SMOS, SMAP, ASCAT, and Copernicus-SSM. These comparisons revealed that $S^2\text{MP}$ consistently outperforms other products in terms of accuracy when benchmarked against ground observations. Furthermore, $S^2\text{MP}$ exhibited strong spatial and temporal consistency with precipitation data from the Global Precipitation Mission (GPM), suggesting that it captures realistic hydrological patterns.

Given the small sample size ($n = 14$), the R^2 was not computed, as it is sensitive to data distribution and may not yield statistically meaningful results under such conditions. Instead, three error metrics were used: bias, RMSE, ubRMSE.

The results of the validation are summarized in Table 5. The $S^2\text{MP}$ product exhibited a slight negative bias of $-0.0180 \text{ m}^3/\text{m}^3$, with an RMSE of $0.0373 \text{ m}^3/\text{m}^3$ and a ubRMSE of $0.0326 \text{ m}^3/\text{m}^3$. In comparison, GSSM-10 showed a small positive bias of $0.0246 \text{ m}^3/\text{m}^3$, an RMSE of $0.0404 \text{ m}^3/\text{m}^3$, and a slightly lower ubRMSE of $0.0320 \text{ m}^3/\text{m}^3$. These findings suggest that both products demonstrate reasonable agreement with ground observations and are capable of capturing surface soil



moisture dynamics with similar levels of accuracy, even under a limited number of validation instances. However, the spatial and temporal applicability of the two products differs significantly. S2MP is limited in geographic scope, covering only selected agricultural plots within parts of France, Morocco, Germany, and the United States. In contrast, GSSM-10 is a globally available product, providing surface soil moisture estimates at 10-meter resolution from 2016 to present. This extensive spatial and temporal coverage makes GSSM-10 more suitable for operational applications in regions where in situ data are sparse and where S2MP is unavailable, thus offering broader utility for global soil moisture monitoring and large-scale environmental assessments.

Table 5. Validation of S2MP and GSSM-10 using 14 in-situ soil moisture observations from USCRN.

Location	Bias (m ³ /m ³)	RMSE (m ³ /m ³)	ubRMSE (m ³ /m ³)
GSSM-10	-0.0180	0.0373	0.0326
SMAP-HB	-0.0246	0.0404	0.0320

4. Applications

4.1 After-fire Assessment

GSSM-10 offers valuable support for wildfire monitoring, assessment, and post-fire recovery planning by capturing high-resolution surface moisture dynamics. For instance, the soil moisture maps of a wildfire-affected region in northern Los Angeles demonstrate a clear contrast between November 10, 2024 and March 22, 2025. The pre-fire condition is illustrated in the Sentinel-2 true color image (Figure 5a) and the corresponding soil moisture map (Figure 5b). The post-fire condition, captured after the Eaton Fire that ignited on January 7, 2025, is shown in the true color image (Figure 5c) and soil moisture map (Figure 5d). The fire burned approximately 14,000 acres in the forested foothills of Los Angeles County, causing extensive environmental damage. The post-fire maps exhibit extensive areas of low soil moisture (reddish-brown tones), highlighting a widespread and persistent surface dryness following the wildfire disturbance. Such depletion of soil moisture can exacerbate erosion risks, delay vegetation recovery, and signal lasting ecosystem stress, especially on steep, fire-exposed slopes prone to debris flows. GSSM-10 can also be used for post-fire impact analysis, restoration monitoring, and land management decisions, providing critical support for assessing ecosystem recovery and mitigating secondary hazards in fire-prone regions.

4.2 Agriculture

The 10-meter resolution of GSSM-10 offers significant utility for irrigated agriculture,

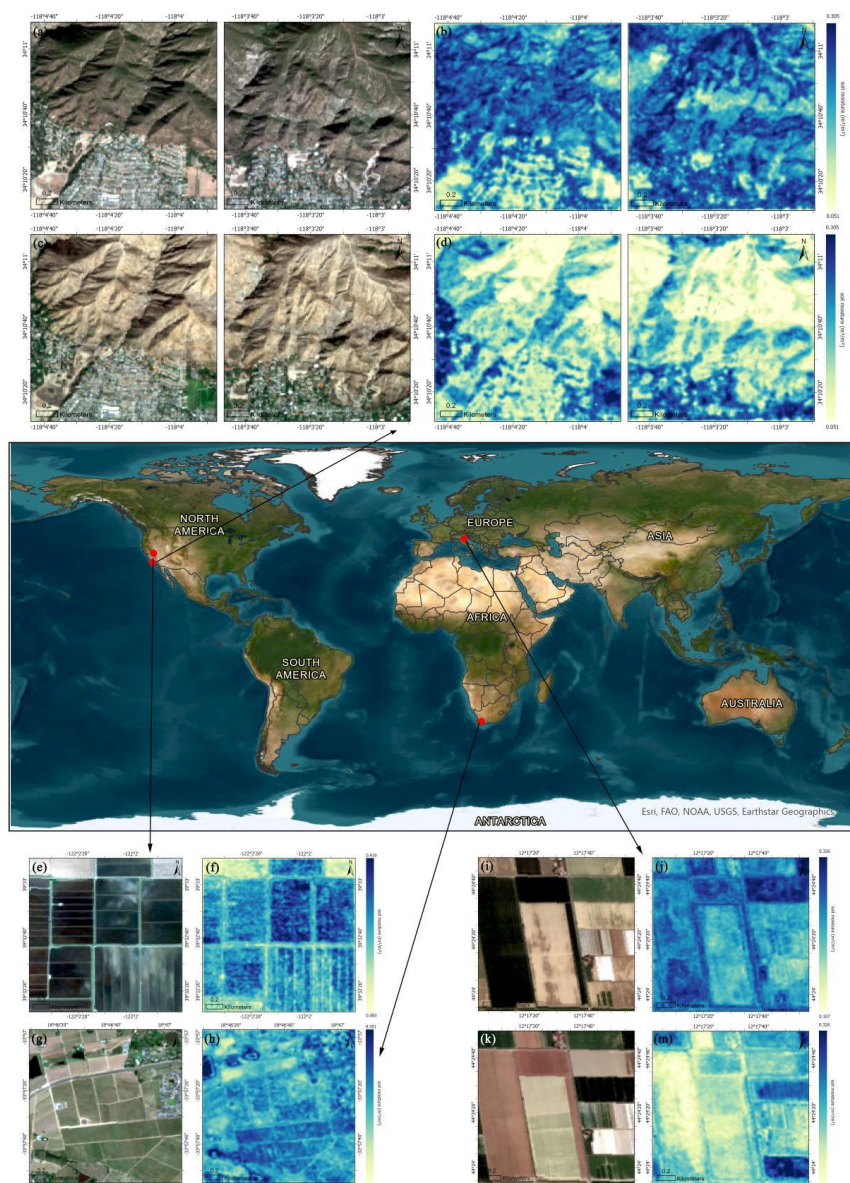


enabling spatially explicit assessments of field-scale soil moisture conditions. This supports more efficient irrigation management, early detection of crop water stress, and optimized resource allocation. Figure 5 presents two representative case studies in agricultural regions. The first site, located in California's Central Valley, is shown in the Sentinel-2 true color image (Figure 5e) and the corresponding soil moisture map (Figure 5f) from November 3, 2016. The second site, located in South Africa, is depicted in the true color image (Figure 5g) and soil moisture map (Figure 5h) from October 20, 2023. The true color imagery delineates field boundaries and management zones, while the corresponding soil moisture maps reveal considerable intra- and inter-field variability in surface moisture conditions.

This spatial variability reflects differences in irrigation and management practices, soil types, vegetation status, and topography—factors that are difficult to capture using coarse-resolution satellite products. The high level of spatial detail provided by GSSM-10 enables applications such as detecting irrigated areas, planning variable-rate irrigation, targeting fertilizer application, detecting early water stress, and evaluating water conservation strategies.

4.3 Flood monitoring

The 10-meter resolution global soil moisture dataset developed in this study provides fine-scale information for analyzing the hydrological impacts of extreme weather events. Figure 5(i) shows a Sentinel-2 true color image, and Figure 5(j) presents the corresponding 10-m resolution soil moisture map for an agricultural area near Ravenna, Italy, on May 6, 2023, shortly after severe rainfall and regional flooding in the Emilia-Romagna region. The soil moisture map reveals widespread saturation, with values ranging from 0.25 to 0.32 m³/m³, consistent with the persistent rainfall that affected the region in early May. In contrast, Figure 5(k) displays a true color image and Figure 5(m) the corresponding soil moisture map for June 27, 2023, following a regional heatwave. This later image reveals markedly drier soils, reflecting the high atmospheric demand and reduced surface moisture after the extreme heat event, during which daily maximum temperatures exceeded 35 °C, rising 8 - 10 °C above long-term average. These observations highlight the dataset's ability to resolve intra-seasonal hydrological variability, effectively capturing both flood-induced soil saturation and subsequent surface drying.



1
 2 Fig. 5. (a) Sentinel-2 true-color image and (b) corresponding 10-m resolution soil moisture
 3 map of northern forest of Los Angeles on Nov. 10th, 2024 (before Eaton Fire); (c) Sentinel-2
 4 true color image and (d) corresponding 10-m resolution soil moisture map of northern forest
 5 of Los Angeles on Mar. 22nd, 2025 (after Eaton Fire).
 6
 7 (e) Sentinel-2 true color image and (f) corresponding 10-m resolution soil moisture map of an
 8 agricultural field in California's Central Valley on Nov 3rd, 2016.
 9 (g) True color image and (h) corresponding 10-m resolution soil moisture map of irrigated



1 fields in South Africa on Oct 20th, 2023.

2

3 (i) True color image and (j) corresponding 10-m resolution soil moisture map near Ravenna,
 4 Italy on May 6, 2023, shortly after extreme rainfall and regional flooding; (k) True
 5 color image and (m) corresponding 10-m resolution soil moisture map on June 27, 2023,
 6 following a regional heatwave.

7

8 5. Code and Data Availability

9

10 The source code and datasets associated with this research are publicly accessible.
 11 The code repository, titled Global-10-m-Surface-Soil-Moisture-Maps, is available on
 12 GitHub: <https://github.com/RSNuo/Global-10-m-Surface-Soil-Moisture-Maps.git>

13 .

14 In addition, the dataset and code have been archived on Zenodo to ensure long-term
 15 accessibility: <https://doi.org/10.5281/zenodo.16956743> (Xu et al., 2025). The Zenodo
 16 record is published under a Creative Commons Attribution 4.0 International (CC BY
 17 4.0) license.

18

19 Users are encouraged to access the repository and archive to reproduce the results and
 20 apply the models to new geographic regions or temporal periods.

21

22 6. Conclusion

23

24 This study presents the development of a global 10-meter resolution surface soil
 25 moisture (GSSM-10) dataset using a multi-sensor, ensemble machine learning
 26 framework. By integrating active microwave, multispectral, thermal, and geographical
 27 data from Sentinel-1, Sentinel-2, Landsat-8/9, and ALOS DSM, and by leveraging
 28 ensemble learning models including TabNet, Random Forest, and XGBoost, we
 29 produced a high-resolution product that substantially advances spatial detail and
 30 predictive accuracy of global SM mapping.

31

32 Our ensemble model demonstrated superior performance compared to individual
 33 models, achieving an R^2 of 0.8344 and an RMSE of $0.0433 \text{ cm}^3/\text{cm}^3$ during
 34 cross-validation, and an R^2 of 0.8591 and RMSE of $0.0401 \text{ cm}^3/\text{cm}^3$ on the test set.
 35 Feature importance analysis highlighted the key roles of geographical features in
 36 shaping soil moisture patterns. The GSSM-10 dataset was validated against in situ
 37 observations and showed comparable or superior accuracy to S2MP and
 38 SMAP-HydroBlocks. Unlike these regionally limited products, GSSM-10 offers
 39 global coverage, 10-m spatial resolution, and near real-time updates, making it a
 40 robust and scalable tool for global soil moisture monitoring and environmental
 41 applications.

42

43

44 The high spatial resolution of GSSM-10 enables a wide range of applications



previously hindered by the coarseness of existing SM products. We demonstrated the utility of this product in wildfire monitoring, irrigated agriculture, and flood analysis, highlighting its ability to resolve fine-scale hydrological variability associated with both natural and anthropogenic disturbances. The dataset also holds promise for supporting ecosystem restoration, climate resilience planning, and precision water management.

Author contribution

N.X. designed the study, processed the satellite and ground datasets, developed the ensemble machine learning framework, and carried out the experiments. A.D. supervised the project, provided critical feedback on methodology and interpretation, and contributed to writing and revising the manuscript. A.A. supported the design of the modeling strategy, contributed to feature analysis and interpretation, and provided comments and revisions on the manuscript. All authors discussed the results and contributed to the final version of the paper.

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