

State of Wildfires 2024-25

Douglas I. Kelley^{1,*}, Chantelle Burton^{2,*}, Francesca Di Giuseppe^{3,*}, Matthew W. Jones^{4,*}, Maria L. F. Barbosa^{1,5}, Esther Brambleby⁴, Joe R. McNorton³, Zhongwei Liu⁶, Anna S. I. Bradley², Katie Blackford⁷, Eleanor Burke², Andrew Ciavarella², Enza Di Tomaso⁸, Jonathan Eden⁹, Igor José M. Ferreira¹⁰, Lukas Fiedler^{11,12}, Andrew J Hartley², Theodore R. Keeping¹³, Seppe Lampe¹⁴, Anna Lombardi⁸, Guilherme Mataveli^{4,10}, Yuquan Qu¹⁵, Patrícia S. Silva¹⁶, Fiona R. Spuler^{17,18}, Carmen B. Steinmann^{19,20}, Miguel Ángel Torres-Vázquez²¹, Renata Veiga²², Dave van Wees²³, Jakob B. Wessel²⁴, Emily Wright², Bibiana Bilbao^{25,26}, Mathieu Bourbonnais²⁷, ~~Gao Cong~~²⁸ ~~Cong Gao~~²⁸, Carlos M. Di Bella^{29,30}, Kebonye Dintwe³¹, Victoria M. Donovan³², Sarah Harris³³, Elena A. Kukvaskaya³⁴, ~~Kukavskaya~~³⁴, ~~Aya~~ Brigitte N'Dri³⁵, Cristina Santín^{36,37}, Galia Selaya^{38,39}, Johan Sjöström⁴⁰, John T. Abatzoglou⁴¹, Niels Andela²³, Rachel Carmenta⁴², Emilio Chuvieco⁴³, Louis Giglio⁴⁴, Douglas S. Hamilton⁴⁵, Stijn Hantson⁴⁶, Sarah Meier⁴⁷, Mark Parrington⁸, Mojtaba Sadegh⁴⁸, Jesus San-Miguel-Ayanz⁴⁹, Fernando Sedano⁴⁹, Marco Turco⁵⁰, Guido R. van der Werf⁵¹, Sander Veraverbeke^{4,15}, Liana O. Anderson⁵², Hamish Clarke⁵³, Paulo M. Fernandes⁵⁴, Crystal A. Kolden⁵⁵

Affiliations:

¹ Water and Climate Science, UK Centre for Ecology & Hydrology, Wallingford OX10 8BB, U.K.

² Hadley Centre, Met Office, Fitzroy Road, Exeter, UK, EX1 3PB

³ European Centre for Medium-range Weather Forecast, Shinfield Park, Reading RG29AX, United Kingdom

⁴ Tyndall Centre for Climate Change Research, School of Environmental Sciences, University of East Anglia,

Norwich Research Park, Norwich, UK, NR4 7TJ

⁵ Federal University of São Carlos, Rodovia Lauri Simões de Barros, km 12 - SP-189 - Aracaju, Buri - São

Paulo, 18290-000, Brazil

⁶ National Centre for Earth Observation, University of Leicester, Space Park Leicester, 92 Corporation Road,

Space City, Leicester, LE4 5SP, UK

⁷ Department of Life Sciences, Imperial College London, Chemistry Building CHEM062, South Kensington,

London, SW7 2AZ, UK

⁸ European Centre for Medium-Range Weather Forecasts, Robert-Schuman-Platz 3, 53175 Bonn, Germany

⁹ Centre for Agroecology, Water and Resilience, Coventry University, Wolston Ln, Ryton-on-Dunsmore, Coventry

CV8 3LG, UK

¹⁰ Earth Observation and Geoinformatics, National Institute for Space Research (INPE), Astronautas Avenue

1758, São José dos Campos, Brazil, 12227-010

¹¹ Institute of Oceanography, Center for Earth System Research and Sustainability, University of Hamburg,

Bundesstraße 53, 20146 Hamburg, Germany

¹² International Max Planck Research School on Earth System Modelling, Max Planck Institute for Meteorology,

Bundesstraße 53, 20146 Hamburg, Germany

¹³ World Weather Attribution, Centre for Environmental Policy, Imperial College London, Weeks Building, 16-18

Prince's Gardens, London SW7 1NE

¹⁴ Water & Climate, Vrije Universiteit Brussel, Pleinlaan 2, 1050 Brussel, Belgium

¹⁵ Faculty of Science, Vrije Universiteit Amsterdam, De Boelelaan 1100, Amsterdam, Netherlands, 1081 HV

¹⁶ School of the Environment, Yale University, 195 Prospect St, New Haven, CT 06511, USA

¹⁷ Department of Meteorology, University of Reading, Brian Hoskins Building, Whiteknights Road, Earley Gate,

Reading, RG6 6ET, UK

¹⁸ The Alan Turing Institute, British Library, 96 Euston Rd., London NW1 2DB

¹⁹ Institute for Environmental Decisions, Department of Environmental Systems Science, ETH Zurich

²⁰ Federal Office of Meteorology and Climatology MeteoSwiss, Operation Center 1, P.O. Box 257, 8058 Zurich-

Airport, Switzerland

²¹ Environmental Remote Sensing Research Group, Department of Geology, Geography and Environment,

Universidad de Alcalá, Calle Colegios 2, Alcalá de Henares 28801, Spain

²² Laboratory for Environmental Satellite Applications (LASA), Department of Meteorology, Federal University of

Rio de Janeiro

²³ BeZero Carbon Ltd, 25 Christopher Street, London, EC2A 2BS, UK

²⁴ Department of Mathematics and Statistics, University of Exeter, Harrison Building, University of Exeter, North

Park Road, EX4 4QF Exeter, United Kingdom

²⁵ Biology, Departamento de Estudios Ambientales, Universidad Simón Bolívar, Valle de Sartenejas, Aartado

89000, Caracas, Venezuela

²⁶ UMR Art-Dev 5281, Université Paul Valéry Montpellier, Site Saint-Charles, France , UMR 5281, Site Saint-

Charles 1, Rue du Professeur Henri Serre, F-34090 Montpellier, France

²⁷ Earth, Environmental and Geographic Sciences, University of British Columbia - Okanagan, 1177 University Way, Kelowna, BC, Canada V1V 1V7
²⁸ Department of Geography, University of Hong Kong, 10F, The Jockey Club Tower, Centennial Campus, Pokfulam Road, Hong Kong
²⁹ Departamento de Métodos Cuantitativos y Sistemas de Información, University of Buenos Aires, Av. San Martín 4453 (1417), CABA, Argentina
³⁰ IFEVA (Agricultural Physiology and Ecology Research Institute), Av. San Martín 4453 (1417), CABA, Argentina
³¹ Department of Environmental Science, University of Botswana, Plot 4775 Notwane Rd, Gaborone Botswana
³² West Florida Research and Education Center, School of Forest, Fisheries, and Geomatics Sciences, University of Florida, 5988 Highway 90, Milton, FL 32583
³³ Fire Risk, Research and Community Preparedness, Country Fire Authority, Burwood East, Victoria, Australia
³⁴ Laboratory of Experimental and Applied Ecology, V.N. Sukachev Institute of Forest Siberian Branch of the Russian Academy of Sciences - separate subdivision of the FRC KSC SB RAS, 50/28 Akademgorodok, Krasnoyarsk, Russian Federation, 660036
³⁵ [Department of Natural Sciences](#), Nangui Abrogoua University, 02 BP 801 Abidjan 02, [Côte d'Ivoire](#)
³⁶ Research Institute of Biodiversity (IMIB), CSIC-University of Oviedo-Principality of Asturias, IMIB, Research Building, Mieres Campus, Mieres, Spain, 33600
³⁷ Biosciences, Swansea University, Wallace Building, Singleton Campus, Swansea, UK, SA2 8PP
³⁸ Research and action, ECOSCONSULT, Calle Tte. H. Balcazar No. 24, Santa Cruz de la Sierra
³⁹ Research, Fundacion Innova, Calle Tte. H. Balcazar No. 24, Santa Cruz de la Sierra
⁴⁰ Fire and Safety, RISE Research institutes of Sweden, Box 857, 51515 Borås, Sweden
⁴¹ School of Engineering, University of California, Merced, 5200 N Lake Rd, Merced, CA, 95343, USA
⁴² Tyndall Centre for Climate Change Research, School of Global Development, University of East Anglia, Norwich Research Park, Norwich, UK, NR4 7TJ
⁴³ Environmental Remote Sensing Research Group, Department of Geology, Geography and Environment, Universidad de Alcalá, Calle Colegios 2, Alcalá de Henares 28801, Spain
⁴⁴ Department of Geographical Sciences, University of Maryland, College Park, MD 20742
⁴⁵ Marine, Earth, and Atmospheric Science, North Carolina State University, 2800 Faucette Drive, Raleigh, North Carolina, USA, 27603
⁴⁶ Program in Earth System Sciences, Faculty of Natural Sciences, Universidad del Rosario, Bogotá, Colombia.
⁴⁷ Land, Environment, Economics and Policy Institute (LEEP), Department of Economics, University of Exeter, Rennes Drive EX4 4ST Exeter, United Kingdom
⁴⁸ Department of Civil Engineering, Boise State University, Boise, ID, USA
⁴⁹ Disaster Risk Management Unit (E.1), Directorate E (Space, Security, and Migration), European Commission Joint Research Centre, European Commission, Rue du Champ de Mars 21, 1050 Brussels, Belgium
⁵⁰ Regional Atmospheric Modelling (MAR) Group, Regional Campus of International Excellence Campus Mare Nostrum (CEIR), Department of Geography, University of Hong Kong
⁵¹ Meteorology and Air Quality, Wageningen University and Research, Droevendaalsesteeg 3-3 A, 6708 PB, Wageningen, the Netherlands
⁵² Cemaden/MCTI, 500 – Distrito de Eugênio de Melo, São José dos Campos – São Paulo, Brazil
⁵³ FLARE Wildfire Research, School of Agriculture, Food and Ecosystem Sciences, University of Melbourne, Grattan St, Parkville, 3010, Australia
⁵⁴ ForestWISE—Collaborative Laboratory for Integrated Forest & Fire Management, Centre for the Research and Technology of Agro-Environmental and Biological Sciences, Universidade de Trás-os-Montes e Alto Douro, Quinta de Prados, Vila Real, Portugal, 5000-801
⁵⁵ Wildfire Resilience Center, School of Engineering, University of California, Merced, 5200 N Lake Rd, Merced, CA 95343, USA

*These authors contributed equally to this work.

Correspondence to:

doukel@ceh.ac.uk
chantelle.burton@metoffice.gov.uk
Francesca.DiGiuseppe@ecmwf.int
matthew.w.jones@uea.ac.uk

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● Abstract

Climate change is increasing the frequency and intensity of extreme wildfires globally, yet our understanding of these high-impact events remains uneven and shaped by media attention and regional research biases. The State of Wildfire Project systematically tracks ~~and analyses~~ global ~~and regional~~ fire activity ~~of each annual fire season, analyses the causes of prominent extreme wildfire events, and this projects the likelihood of similar events occurring in future climate scenarios~~. This, its second annual report, covers the March 2024 to February 2025 fire season. During the 2024-25 fire season, fire-related carbon (C) emissions were totalled 2.2 Pg C, 9% above average and the 6th highest on record since 2003, despite below-average global burned area (BA; ~~3.7 million km²~~). Extreme fire seasons in South America's rainforests, dry forests and wetlands, and in Canada's boreal forests pushed up the global C emissions total. Fire C emissions were over four times above average in Bolivia, three times above average in Canada, and ~50% above average in Brazil and Venezuela. Wildfires in 2024-25 caused 100 fatalities in Nepal, 34 in South Africa, and ~~3031~~ in Los Angeles, with additional fatalities reported in Canada, Côte d'Ivoire, Portugal, and ~~Turkey~~~~Türkiye~~. The Eaton and Palisades fires in Southern California caused 150,000 evacuations and US\$140 billion in damages. Communities in Brazil, Bolivia, Southern California, and Northern India were exposed to fine particulate matter at concentrations 13-60 times WHO's daily air quality standards. We evaluated the causes and predictability of four extreme wildfire episodes from the 2024-25 fire season, including in Northeast Amazonia (January-March 2024), the Pantanal-Chiquitano border regions of Brazil and Bolivia (August-September 2024), Southern California (January 2025), and the Congo Basin (July-August 2024). Anomalous weather created conditions for these regional extremes, while fuel availability and human ignitions shaped spatial patterns and temporal fire dynamics. In the three tropical regions, prolonged drought was the dominant fire enabler, whereas in California, extreme heat, wind, and antecedent fuel build-up were ~~the dominant~~~~compounding~~ enablers. Our attribution analyses show that climate change made extreme fire weather in Northeast Amazonia 30–70 times more likely, increasing ~~burned area~~~~BA~~ roughly fourfold compared to a scenario without climate change. In the Pantanal-Chiquitano, fire weather was 4–5 times more likely, with ~~up to~~ 35-fold increases in ~~burned area~~~~In BA~~. Meanwhile, our analyses suggest that BA was 25 times higher in Southern California, ~~due to climate change made larger burned area 89% more likely, with burned area up to 25 times higher~~. The Congo Basin's fire weather was 3–8 times more likely with climate change, with a 2.7-fold increase in ~~burned area~~~~BA~~. Socioeconomic changes since the pre-industrial period, including land-use change, also likely increased ~~burned area~~~~BA~~ in Northeast Amazonia. Our models project that events on the scale of 2024-25 will become up to 57%, 34%, and 50% more frequent than in the modern era in Northeast Amazonia, the Pantanal-Chiquitano, and the Congo Basin, respectively, under a ~~middle-of-the-road~~ ~~medium-high~~ scenario (SSP370-) by 2100. Climate action can limit the added risk, with frequency increases ~~kept~~~~held~~ to below 15% in all three regions under a strong mitigation scenario (SSP126). In Southern California, the future trajectory of extreme fire likelihood remains highly uncertain due to poorly constrained climate-vegetation-fire interactions influencing fuel moisture, though our models suggest that risk may decline in future. This annual report from the State of Wildfires Project integrates and advances cutting-edge fire observations and modelling with regional expertise to track changing global wildfire hazard, guiding policy and practice towards improved preparedness, mitigation, adaptation, and societal benefit. Thirteen new datasets and model ~~codebase's~~~~codebases~~ presented in this work are available from the State of Wildfires Project's Zenodo community (<https://zenodo.org/communities/stateofwildfiresproject>, last access: 11 August 2025).

● Short Summary

The second State of Wildfires report examines extreme wildfire events from 2024 to early 2025. It analyses key regional events in Southern California, Northeast Amazonia, Pantanal-Chiquitano, and the Congo Basin, assessing their drivers, predictability, ~~and attributing them~~

the influence of climate change and land use. Seasonal outlooks and decadal projections are provided. Climate change greatly increased the likelihood of these fires, and without strong mitigation, such events will become more frequent.

1. Introduction

1.1. Background

The potential for wildfires is growing under climate change, with increases in the frequency and intensity of drought and periods of fire-favourable weather driving reductions in vegetation (fuel) moisture and priming landscapes to burn more regularly, intensely, and severely (Seneviratne et al., 2022; UNEP, 2022a; Jones et al., 2022; Abatzoglou et al., 2019; Cunningham et al., 2024a). Additionally, human activities and land use change can contribute to or exacerbate the risk of extremely large, fast-moving or intense fires, especially in tropical forests where people are the primary cause of ignition and forest degradation (Lapola et al., 2023). Recent years have been marked by a series of extreme wildfire events spanning the globe (Abatzoglou et al., 2025), with record levels of burned area (BA) occurring in 2019–2020 Australian “Black Summer” bushfires (Abram et al., 2021; Canadell et al., 2021) and a series of, while successive high-ranking wildfire seasons occurring in quick succession took place in the western US (2020 and 2021; Higuera & Abatzoglou, 2020), Siberia (2020 and 2021; Zheng et al., 2023), varying parts of Europe (e.g. 2017, 2022, 2023; European Commission Joint Research Centre, 2023, 2024, 2025), South America (2019, 2020, 2023, 2024; Kelley et al., 2021; Barbosa et al., 2022; Silveira et al., 2020; Mataveli et al., 2024, 2025), and Canada (2023, 2024; Jones et al., 2024b; Jain et al., 2024; Byrne et al., 2024; Kolden et al., 2024). The 2024–25 fire season was marked by extreme fire extent and emissions in Amazonia and the Pantanal-Chiquitano (Mataveli et al., 2025; Kolden et al., 2025) and a second consecutive year of extreme fire extent and emissions in Canada (Kolden et al., 2025; Parrington and Di Tomaso, 2025). The 2024–25 fire season also saw extreme Equatorial Africa experienced severe fire activity in equatorial Africa, which went broadly under-reported that received little international attention despite fires triggering driving record rates levels of forest loss (stand-replacing fire extent) in the region (World Resources Institute, 2025). Meanwhile, extremely In addition, highly destructive and costly individual fires affected Southern California (Barnes et al., 2025; Woolcott, 2025) and Jasper National Park in Alberta (Parks Canada, 2024; Insurance Bureau of Canada, 2025). Widespread regional anomalies in high fire activity were also seen recorded in northern India leading and Nepal, where high fire activity contributed to severe haze event episodes (CAMS, 2024) and at least 100 people were killed (Bolakhe, 2024).

The prominence of recent extreme wildfires and wildfire seasons notably contrasts with overall trends in the area burned by fires globally. A distinctive These recent high-impact events align with a growing trend has emerged towards enhanced fire activity increasing extent, intensity and severity in forests and other fuel-rich environments, which is occurring amidst increasingly frequent and intense droughts and heatwaves of fires globally, particularly in the extratropics. Key trends observed over approximately the past two decades include a ~40% increase in the extent of forest fires, predominantly in the extratropics, and a 2-fold increase in the intensity of globally extreme fires (Jones et al., 2024a; Cunningham et al., 2024a). Due mostly These changes are linked to both climatic and human factors. Climate change is intensifying the frequency and severity of drought and fire-favourable weather conditions, lowering vegetation (fuel) moisture and preconditioning landscapes to burn more regularly, intensely, and severely (Seneviratne et al., 2022; UNEP, 2022a; Jones et al., 2022; Abatzoglou et al., 2019; Cunningham et al., 2024a). a reduction in the At the same time, human activities, including land use and land use change, exacerbate the risk of large, fast-moving, or intense fires, especially in tropical forests where they remain the dominant cause of ignition and degradation (Lapola et al., 2023; World Resources Institute, 2025).

The prominence of recent extreme wildfire events contrasts with the broader global savannahs trend in BA. Since the early 2000s, global BA has declined by roughly one-quarter, driven mainly by reductions in the area burned by fires in African savannahs linked to landscape fragmentation and changing changes in the distribution of rainfall patterns, global BA has fallen since the beginning of this century by around one-quarter (Andela et al., 2017; Zubkova et al., 2019; Jones et al., 2022; Chen et al., 2024). Critically, this decline in fire extent masks major shifts in the distribution of fires globally, with regions such as eastern Siberia and the western US and Canada experiencing a more than 40% increase in BA since 2000 (Jones et al., 2022; Zheng et al., 2021) and regions such as southeast Australia also showing significant increases over longer periods despite high interannual variability (Canadell et al., 2021). Likewise, there have been shifts in the global distribution of BA from non-forests to forests globally and from the tropics to the extratropics, with the increased prevalence and severity of forest fires emitting increasing quantities of forest carbon stocks each year and driving increasing fire carbon (C) emissions globally (Kolley et al., 2019; Jones et al., 2024a). Hence, focussing exclusively on global aggregated BA extent underplays the scale and magnitude of the significant shifts in wildfire activity and impacts that are underway across many world regions. An increase in forest and peatland burning is particularly concerning due to the rich ecosystem services that these regions provide, including C storage and biological and cultural diversity (UNEP, 2022b). The intensification of fire regimes in environments that are less fire-adapted is of high importance because these ecosystems are expected to be least resilient to such changes (Grau-Andrés et al., 2024) and because they are often home to communities relying directly on the forest. Excessive focus on global aggregated BA totals, which are principally responding to trends in low-intensity fire regimes in savannahs, underplays the scale and magnitude of the significant shifts in wildfire activity and impact that are underway across many world regions. forest (Newton et al., 2022; Shepherd et al., 2020; Schleicher et al., 2018).

The extreme wildfire events of recent years have significantly impacted brought significant impacts on societies and ecosystems across the globe (Cunningham et al., 2024a, 2024b; 2025). Since 1990, wildfire disasters have directly killed or injured at least ~18,000 people, a conservative measure based on incomplete records and reporting biased to the global Northern countries (updated from Jones et al., 2022; Centre for Research on the Epidemiology of Disasters, 2024). In 2023, 232,000 people were evacuated due to wildfires in Canada alone (Jain et al., 2024; Kolden et al., 2024). Also since 1990, fires are estimated to have caused on the order of 1.5 million premature deaths globally per year through degraded air quality related to fine particulate matter (PM_{2.5}; Johnston et al., 2021; Xu et al., 2024; Chen et al., 2021). Degraded air quality related to fires is experienced most strongly in the tropics (Pai et al., 2022) and often disproportionately affects the elderly, the young, the infirm, and traditional communities with poor public services or means of protection (Johnston et al., 2021; Carmenta et al., 2021). Extreme fires can moreover impact the livelihoods of various communities and landowners who depend on intact natural landscapes. For example, the lands, territories and cultural heritage of traditional communities and Indigenous Peoples can be degraded and transformed by wildfires, raising climate justice issues that compound a legacy of colonisation, dispossession and forced cessation of cultural practices (Garnett et al., 2018; Barlow et al., 2018; Lapola et al., 2023; Pascoe et al., 2024 Carmenta et al., 2021; Johnston et al., 2021).

Extreme fires are also a threat to ecosystem services, including biological and cultural diversity and C storage (Jones et al., 2024; UNEP, 2022b). The intensification of fire regimes in environments that are less fire-adapted is particularly critical because these ecosystems are expected to be least resilient to such changes (Grau-Andrés et al., 2024) and because they are often home to communities relying directly on intact forest (Newton et al., 2022; Shepherd et al., 2020; Schleicher et al., 2018). As anthropogenic emissions of CO₂ remain persistently high, the world's natural C sinks in forests, peatlands, marine and other ecosystems are increasingly pivotal to moderating increases in atmospheric CO₂ concentration emissions (Friedlingstein et al., 2025). Intact forests are often relied upon for

delivering national plans for reaching Net Zero (Smith et al., 2023) and offering sites for nature based solutions (NBS). Yet, massive wildfire emissions from boreal forests and soils in Siberia and Canada amount to over 1 billion tonnes of C across the years 2020, 2021, and 2023 ~~amount to over 1 billion tonnes of C alone~~, a gross flux comparable in magnitude to annual CO₂ emissions from fossil fuel combustion in India, the EU27, or the USA (Friedlingstein et al., 2025; Zheng et al., 2023). In a natural fire regime, these gross emissions would likely be recuperated through post-fire recovery. However, ~~the greater vegetation mortality and loss of ecosystem function associated with shifts in fire regime towards~~ more widespread and severe fires ~~contribute have contributed~~ to shifts in local to regional a reversal of terrestrial carbon budgets from sinks to sources in some regions, driven by the enhanced disturbance of vegetation and soils carbon stores (Zheng et al., 2021; Gatti et al., 2021; Nolan et al., 2021a; Phillips et al., 2022; Harrison et al., 2018; Jones et al., 2024a; Burton et al. 2024). Loss of vegetation during extreme fire seasons can also have wider lasting effects on ecosystems, for instance by reducing the habitat area available to endemic species (Ward et al., 2020; Carmenta et al., 2025). Emerging evidence further suggests that extreme fire events can influence marine carbon sinks, through deposition of pyrogenic aerosols and nutrients that alter ocean biogeochemistry, creating feedbacks between terrestrial fires and the global carbon cycle (Tang et al. 2021)

~~Extreme fires can moreover impact the livelihoods of various communities and landowners who depend on intact natural landscapes. For example, the lands, territories and cultural heritage of traditional communities and Indigenous Peoples can be degraded and transformed by wildfires, raising climate justice issues that compound a legacy of colonisation, dispossession and forced cessation of cultural practices (Carnott et al., 2019; Barlow et al., 2019; Lapola et al., 2023; Pascoe et al., 2024). Further, conflating the detrimental impacts of wildfire types has also stigmatised small scale intergenerational fire use and led to prohibitive fire governance that affects local communities (Carmenta et al., 2021; Barlow et al., 2020; Pascoe et al., 2024).~~

Mitigating and adapting to increases in wildfire potential are growing priorities of policymakers and require coordination with many other stakeholders. National and international disaster management centres are seeking to enhance predictive capacity, while fire management agencies are expanding or re-allocating their resources to rapidly suppress fires to avoid them becoming too large, fast, or intense (e.g. Bowman et al., 2020). ~~A number of various~~ international organisations such as the UN Environment Programme (UNEP, 2022a), the World Bank (2020, 2024a), the Organisation for Economic Co-operation and Development (OECD, 2023), and a range of other inter- or non- governmental organisations are producing reports that consolidate evidence on the changing risk of extreme fires and identify best practices for mitigating their impacts, including through land management ~~and~~, urban/rural planning, ignition risk reduction, pre-suppression, rapid early/initial attack. Many land managers are developing and implementing approaches such as fuel reduction (Fernandes and Botelho, 2003; Stephens et al., 2012; Moreira et al., 2020; Chuvieco et al., 2023; Hsu et al., 2025). Wildfire response agencies are exploring innovative approaches to detecting and responding to fires, and there is rising interest in the prospect of integrated fire management around the world (Food and Agriculture Organization of the United Nations, 2024). Operators of C market projects and forest carbon-conservation initiatives, such as REDD+ are particularly wary of the risks that wildfires present to the permanence of C offsets, which often feature as a key tool in national policies and international initiatives for achieving Net Zero emissions (Barlow et al., 2012; Smith et al., 2023).

Not all landscape fires are 'bad fires'. Many ecosystems are fire-adapted, with flora that have developed competitive advantages to defend against damage from fire or to resprout or regenerate after fire, and fauna that exploit the habitats created by fire-adapted vegetation (Kelly et al., 2020; Pausas and Keeley, 2023). As Pausas et al. (2025) note, fire is a "defining feature of our biosphere, having appeared when the first plants colonized the land, and it

continues to occur across the planet at different frequencies and intensities". In addition, fire has played a vital role in the success of the human species, from its early domestication for cooking, warmth and protection, through millennia of cultural burning to shape landscapes and resources (Bowman et al., 2011; Pyne, 2011). Small-scale intergenerational fire use continues to be used by Indigenous and traditional communities around the world, and to label all fire as 'bad fire' would risk erasing culturally embedded stewardship, stigmatising traditional practices and cultural values, undermining livelihoods and biodiversity, and increasing future wildfire risk by preventing the low-intensity cultural burns that maintain habitat mosaics and keep hazardous fuels in check (Carmenta et al., 2021; Barlow et al., 2020; Pascoe et al., 2024). The practice of low-intensity prescribed burning, which recognises the need for fire on fire-adapted landscapes, is applied in many world regions for the purpose of hazardous fuel reduction or for the rejuvenation of vegetation aligned with vegetation adaptations, often with inspiration from cultural burning practices (Hiers et al., 2020; Hsu et al., 2025). Nonetheless, trends towards larger, more intense or more severe fire properties have the potential to push fire-adapted ecosystems towards the edge of their physiological range (Kelly et al., 2020; Pausas et al., 2025). At the same time, low-intensity controlled burning are, in some regions, facing shrinking windows of weather conditions in which low-intentional burns can be safely maintained (Fernandes et al., 2013; Swain et al., 2023; Di Virgilio et al., 2020).

Amidst extreme wildfires and wildfire seasons, stakeholders increasingly turn to scientists for answers. How extreme was this fire event in a historical context? Is climate change amplifying fire occurrence? Can we disentangle the factors responsible in order to target those in policy and management? Will we see more wildfires like this in the future? Did land use or management factors exacerbate or ameliorate the problem? Could we have predicted these events and how can we improve early warning systems and preparedness in the future? What is the role of climate and socioeconomic factors, such as land use, in reducing risk of extreme wildfires in future?

While observational, statistical, and modelling tools for assessing extreme wildfire drivers and predicting wildfire occurrence are advancing rapidly, their application to studying extreme wildfire seasons or events on timescales relevant to public and political interest remains limited. The State of Wildfires report represents ~~a new initiative~~ an annual effort to systematically catalogue extreme wildfire events at annual frequency and explain their occurrence, predictability and attribution to climate and land use changes. As well as collating records of extreme fire activity from Earth observations, we convened regional expert panels for each continent to help identify events considered extreme in terms of their social, economic, and ecological impacts, thereby capturing important dimensions of fire activity not always visible in satellite data. The report incorporates recent methodological advances in disentangling the drivers of four selected extreme wildfire events to fuel dryness, fuel load, and weather, and ignition and suppression factors. By applying these methodological advances in conjunction with models of global change, we quantify the change in likelihood of the past year's events under climate and land use changes. Observable fire metrics (e.g. BA) are the target variable of our causal inference and attribution work, which thereby advances on more common climate attribution studies that attribute change in fire-favourable meteorological conditions to climate change. Overall, this report capitalises on recent advances in the study of extreme fire events and seasons to provide timely information about shifting fire regimes and their causes. The findings of the report are relevant to organisations involved in wildfire prevention and combat efforts, policymakers, the media, and the wider public.

1.2. Objectives of this Report

~~The State of Wildfires report aims to deliver actionable information to policy and practice stakeholders and wider society. In rising to this challenge, we aim to spur scientific and~~

technological innovation including stimulating development of better tools for understanding and predicting extreme fires. In this edition we:

The *State of Wildfires* report brings together the latest science on extreme fire monitoring, prediction, and modelling to track how wildfire impacts on society and the environment are changing, and to explain the drivers behind these shifts. It forms the foundation of the wider *State of Wildfires Project*, which aims to deliver insights on climate, land use, and fire management policy to decision-makers and practitioners, ultimately supporting stronger societal and environmental resilience to wildfire. In this edition of the *State of Wildfires Report* we:

1. Regionally identify extreme individual wildfires or extreme wildfire seasons of the period March 2004-February 2025, and place them in the context of recent trends- (Section 2).
2. ~~Quantify the impacts of extreme~~Shortlist four 'focal' extremes events (extreme individual wildfires or extreme wildfire seasons) with notable impacts on society or the environment, which are the subject of dedicated analysis in subsequent report sections (Section 2).
- 2-3. Globally assess the impacts of extreme wildfire events in terms of the exposure of population, physical assets (built environment), and carbon projects to fire as well as ~~degraded~~their impacts on air quality- (Section 3).
3. ~~Shortlist a selection of four extremes (extreme individual wildfires or extreme wildfire seasons) with notable impacts on society or the environment, which we term the 'focal events'.~~
4. Diagnose the contributions of ~~both~~weather, fuel dryness ~~and~~, fuel load, ignitions, and suppression to the occurrence of each focal event- (Section 4).
5. Assess the capacity of operational predictive systems to predict the scale of fire occurrence in each focal event- (Section 4).
6. Attribute each focal event to anthropogenic influences by testing the role of climate change and socioeconomic factors such as land use, land use change, and human ignitions- (Section 5).
7. Provide an outlook on the probability of extreme events in the ~~coming early months of the 2025-26~~ fire season (~~commencing March 2025~~Section 6).
8. Project future changes in the probability of each focal event under future climate scenarios- (Section 6).

~~Key methodologies used to achieve the above objectives are summarised as follows. The delivery of this report and its objectives relies on critical datasets and models developed by the research community over many decades. By applying these tools to the challenge of studying extreme wildfires, we not only gain insights into their strengths and limitations for studying extreme fires, but also drive scientific and technological innovation that advances our ability to monitor, explain, and predict such events.~~ To address objectives 1 and 2, we build a comprehensive dataset of fire metrics including BA, fire counts, fire C emissions, and regional statistics of individual fire properties ~~(such as~~ size and rate of growth), all for consistent world regions and quantitatively identify anomalies in these metrics during the past fire season (Giglio et al., 2018; van der Werf et al., 2017; Andela et al., 2019). To address objectives ~~34~~ and ~~45~~, we leverage weather forecasts from the European Centre for Medium-Range Weather Forecasts (ECMWF) at different time horizon from medium (1-15 days) to long range (up to 4 months ahead) and additionally employ two state-of-the-art fire models, *Controlar Fogo Local Analise pela Máxima Entropia* - English "Local Fire Control Analysis by Maximum Entropy" (ConFLAME; Kelley et al., 2019; Barbosa et al., 2025b) and *Probability of Fire* (PoF; McNorton et al., 2024) to pinpoint the causes of the extreme fire events of 2024-25. To address objective ~~56~~, we employ projections of fire weather from the Hadley Centre Large Ensemble (HadGEM3-A, Ciavarella et al., 2018) to attribute change in the Fire Weather Index (FWI) to

climate change- (Abatzoglou et al., 2019; Clarke et al., 2022), and we drive ConFLAME (Kelley et al., 2019; Barbosa et al., 2025b) with outputs from both HadGEM3-A and separately with the Intersectoral Impacts Model Intercomparison Project 3a (ISIMIP3a) and Joint UK Land Environment Simulator Earth System model (JULES-ES; Mathison et al., 2023) to attribute extreme BA to climate and land use changes (Burton, Lampe et al., 2024). To address objective 67, we use seasonal outlook of FWI from the Copernicus Emergency Management Service (Di Giuseppe et al., 2024). To address objective 78, we again pair ConFLAME with JULES-ES (Mathison et al., 2023) to project future changes in BA under several future climate and land use scenarios and provide a comprehensive assessment of past and future extreme wildfire events.

The State of Wildfires report was launched in 2024 and is an annual report that can harness and adopt new methodologies brought forward by the scientific community in the interim between its yearly publication. Over the coming years and decades, we aim to enhance the tools presented in this report to predict extremes with increasing lead times, monitor emerging situations in near-real time, and explain their causes rapidly, thus enhancing our ability to deliver timely insights to decision-makers when they are most needed.

2. Extreme Wildfire Events of 2024-25

2.1. Methods

We catalogued the extreme regional wildfire events or annual fire seasons in the period March 2024-February 2025 based on a combination of anomalies in the distribution of several observable fire metrics from Earth observations (Section 2.1.1 and Section 2.1.2). In this work, the global fire season is defined as occurring in March-February windows oriented around the annual minima of global fire activity in boreal spring (see further details in Section 2.1.1.2). As a new development for this edition of the report, we added statistics describing anomalies in fire intensity (with respect to fire radiative power; see section 2.1.1) during the 2024-25 fire season, complementing anomaly statistics provided in the prior edition related to regional BA, fire emissions, fire size, and rate of growth.

Due to the diversity of environmental settings in which fires occur and the range of ecological, economic, or societal impacts caused, defining an extreme fire or an extreme fire season remains inherently challenging. To date, extreme fires have commonly been defined by their BA extent, by their feedback on the global climate, and by their socio-economic and ecological impacts (Linley et al., 2022, 2025; Driscoll et al., 2024). We reviewed the range of approaches that can be taken to identify extreme wildfire events in our inaugural report (see Appendix A of Jones et al., 2024b) and so do not revisit this in the current article B1.1 of Jones et al., 2024b). A universally accepted objective definition of “extreme” remains elusive, reflecting a series of data- and knowledge-oriented challenges. Data-oriented challenges include the absence of consensus on quantitative criteria, with no universally applicable thresholds for size, severity, or other measurable properties; pronounced geographic variability, as regional fire regimes dictate the relevance of particular thresholds; evolving definitions that have progressively expanded to encompass a broader range of fire types and behaviours under changing climatic conditions; and context dependence, whereby interpretation is contingent on ecosystem characteristics, historical fire regimes, and benchmarks such as return intervals or ecosystem damage. Knowledge-oriented challenges centre on the lack of agreement over qualitative criteria, including fire behaviour and impacts; the proliferation of overlapping and sometimes redundant terminology (e.g. “megafire,” “catastrophic fire”); the influence of linguistic and cultural context on interpretation and reporting; the shaping of scientific terminology by societal discourse, necessitating accessibility to diverse audiences; and the limited rigour, clarity, and standardisation evident in existing definitions. Recognising these complexities and the need for transdisciplinary processes to establish robust, standardised

criteria in future work, this report maintains a deliberately broad and flexible definition of extreme.

While an extreme fire event or extreme fire season may be visible as a significant anomaly against historical Earth observations, the scientific community seeks to apply a more comprehensive definition of extreme fire, including its impacts on society and the environment. To catalogue extreme events that were not necessarily visible in Earth observations, regional expert panels were constructed and given responsibility for identifying extreme events of the past fire season (**Section 2.1.3**). The expert panels were given flexibility to identify and catalogue wildfire characteristics or impacts that are considered regionally extreme but are not necessarily captured by Earth observations. Examples of extremes that can be captured by expert assessment (but not by Earth observations) include: suppression difficulty; fatalities and structure loss; impacts on human health and wellbeing; impacts on agricultural and other economic sectors; impacts on biodiversity; and; impacts on diverse ecosystem services such as recreation, tourism, or other cultural values. extremes that are not captured by Earth observations, such as suppression difficulty or impacts on society or the environment (see **Section 2.1.3**). Hence, **Section 2.2** identifies a variety of impactful events displaying a broad range of characteristics and impacts that can occur across diverse fire regimes (e.g. Archibald et al., 2009; Cunningham et al., 2024a, 2024b; Keeley, 2009).

As a new development for this report, we added several new analyses providing context to the observed extremes in fire during the past fire season (**Section 2.1.4**). Specifically, we added an analysis of extreme (95th percentile) fire weather days for the 2024-2025 fire season allowing the spatial and temporal context of extreme fires with extreme fire weather to be described. (**Section 2.1.4.1**).

2.1.1. Earth Observations of Fire

2.1.1.1. Input Datasets

We assembled observations of burned area (BA), synonymous with fire extent, for the period March 2002-February 2025 from the National Aeronautics and Space Administration (NASA) product MCD64A1 (collection 6.1). MCD64A1 provides daily BA observations at 500 m spatial resolution with global coverage and is based on retrievals from the Moderate Resolution Imaging Spectroradiometer (MODIS) sensors mounted to the Terra and Aqua satellites (Giglio et al., 2018, 2021).

We also produced a global record of individual fires for the period March 2002-February 2025 by updating the Global Fire Atlas (Andela et al., 2019) through February 2025, driven by the 500m MODIS BA data. The Global Fire Atlas algorithm clusters burned cells into individual fires, tracks their daily progression, and logs attributes such as fire size and mean daily rate of growth. Our updates are provided at Andela and Jones (2025). The Global Fire Atlas is one of several products tracking daily fire progression and identifying individual fires at global scale based on moderate resolution satellite data (Andela et al., 2019; Laurent et al., 2018; Artés et al., 2019). The product uses the MODIS BA product. The smallest unit of disaggregation is 500m and the shortest timestep on which the expansion of a fire can be observed is daily. Given its resolution, the Global Fire Atlas is expected to represent the dynamics of large fires better than smaller fast-moving fires.

In addition, we gathered estimates of fire carbon (C) emissions for the period March 2024-February 2025 from two models driven by Earth observations of active fires or BA: firstly, the Global Fire Assimilation System (GFAS) product, provided operationally by the Copernicus Atmospheric Services (CAMS) at 0.1 degree spatial resolution (approximately 11km at the equator) and daily temporal resolution (Kaiser et al., 2012; European Centre for Medium-

Range Weather Forecasts, 2024), and; secondly, the Global Fire Emissions Database (GFED; version 4.1s) product at 0.25 degree spatial resolution and daily temporal resolution (van der Werf et al., 2017). GFAS is driven by the fire radiative power (FRP) retrievals in the MODIS active fire product MCD14A1 and biome-level relationships between FRP and biomass consumed based on GFED3 (Kaiser et al., 2012). For the 1997-2016 period, GFED4s is driven by MODIS BA data (MCD64A1 collection 5) supplemented with small fire BA based on MODIS active fire data, and a model for biomass productivity and fuel consumption (van der Werf et al., 2017). For the post-2016 period in GFED4.1s, emissions are based on estimates that rely solely on MODIS active fire data, with pixel-level scaling factors trained on the relationship between active fire detections and burned-area-driven emissions using pixel-based scaling factors derived from the 2003-2016 overlapping period.

As a new analysis developed for the 2024-25 report, we added summaries of the peak (95th percentile) intensity of the fires detected in the Global Fire Atlas. The underlying data for this analysis were daily observations of fire radiative power (FRP) from the NASA active fire products MOD14A1 and MYD14A1 (Giglio et al., 2016). FRP measures the rate of radiant energy emitted by a fire, which is directly related to the fire's intensity and fuel consumption. These products report FRP as the instantaneous rate of radiative energy release from actively burning fires, expressed in megawatts (MW) per pixel. Conceptually, FRP provides a satellite-derived measure of the combustion rate, and is therefore more closely aligned with fire radiative intensity (W m^{-2}) than with measures of fireline intensity (W m^{-1}) used more frequently in field studies, fire behaviour modelling, and active fire management. MOD14A1 and MYD14A1 each provide FRP observations at two different times of the day, with the MOD14A1 dataset produced based on retrievals from the MODIS sensor aboard NASA's Terra satellite, which overpasses at around 10:30 AM and 10:30 PM local time, and the MYD14A1 dataset produced based on retrievals from the MODIS sensor aboard NASA's Aqua satellite, which overpasses at around 1:30 PM and 1:30 AM local time. In our case, daytime and nighttime observations of FRP were combined into a single dataset of active fire detections obtained from any satellite overpass and either MODIS sensor. To minimize potential uncertainties, we excluded FRP measurements associated with large MODIS scan angles ($>50^\circ$), and normalized the FRP measurements by pixel size (Li et al., 2024).

The upcoming decommissioning of the Terra and Aqua satellites on which the MODIS instruments are mounted pose potential challenges for evaluating long-term data records of BA and estimated emissions from wildfires. The wider community requires continued development of BA and active fire products from sensors such as VIIRS (e.g., Parrington et al., 2025).

2.1.1.2. Input Data Uncertainties

We note that the MODIS BA product data used in our analyses of anomalies in BA and individual fire properties (via the Global Fire Atlas) are known to be conservative due to the limitations to detecting small or short-lived fires (e.g. agricultural fires) based on surface spectral changes at 500m resolution. Recent work has shown that including detections of small active fires increases global BA estimates by 93% (Chen et al., 2023). However, variability and trends in regional BA totals using datasets that include small fires do not differ significantly from the variability and trends present in the MODIS BA product (Chen et al., 2023). Hence, inclusion or exclusion of small fires tends to generate biases in central estimates of BA in one direction or the other, in line with the sensitivity of different sensors to different fire types. Uncertainty in the detection of small fires is larger than in the case of fires detected in the MODIS BA product, due to limited validation (van der Werf et al., 2017). The MODIS BA product with resolution of 500 m is deemed highly suitable for addressing the research questions of this report, which focus on more impactful fires that tend to burn larger areas. MCD64A1 BA product because the corrections made for small fires are consistent over

time (Chen et al., 2023). In this report, we require a BA product that has global coverage and ideally a spatial resolution that can be aggregated within geographical divisions (e.g. Table 1), and also a temporal consistency over a multi-decadal time series up to the present year. The MODIS MCD64A1 BA product meets these needs and allows us to address our research questions, though we caution that the absolute values of BA reported by the MODIS BA product are underestimated due to small fire omission.

Uncertainties in the BA estimation can be approached by comparing different existing global BA products. For instance the estimations of BA from the NASA MCD64A1 product, which is the basis for the calculations of this paper are 40% lower than ESA FireCCI51, based on Sentinel-3 reflectance and VIIRS active fires, and 20% lower than the estimations provided by the Copernicus Land service (March 2024-Feb 2025 period). Comparing these estimates with the BA derived from higher resolution sensors, such as Sentinel-2 MSI would probably double the estimations of MCD64A1, as it was observed in Africa (Chuvieco et al., 2022) and the GFED5 BA product (Chen et al., 2024).

Uncertainties in the BA estimation can be decomposed into observational uncertainty and parameter uncertainty. Observational uncertainty arises due to errors of omission (missed detections of true fires) and commission (incorrect identification of fires that did not occur). The NASA MODIS MCD64A1 was previously assessed to have a 40% commission error and 73% omission error, with an overall 50% under-detection of BA versus higher resolution Landsat imagery due to omitted small fires (based on analysis from 558 sites selected by probability sampling). Note that these uncertainties are highly consistent with the reported 93% increase in BA in the GFED5 BA product versus MCD64A1, which occurs due to small fire omissions in MCD64A1 (Chen et al., 2024). Other work comparing the ESA FireCCI51 BA product (based on 250 m MODIS retrievals) with the high resolution ESA FireCCI51 BA product (based on 20m Sentinel-2 retrievals) also suggests an approximately 50% under-detection of BA in sub-saharan Africa (Chuvieco et al., 2022). Parameter uncertainty arises from the range of methodological choices that can be made when producing a global BA product based on the same or similar observational inputs. These uncertainties can be quantified by comparing different global BA products using similar observations as input. For instance, the estimates of global BA from the NASA MODIS MCD64A1 product (based on 500m MODIS observations) are 20% lower than the BA estimates provided by the Copernicus Land service (2025; based on 300 m Sentinel-3 observations with similar sensor properties and with no corrections for small fires as in ESA FireCCI products).

Uncertainties in fire carbon emissions estimates from GFED4.1s are on the order of $\pm 20\text{-}25\%$ at 1 standard deviation for global totals (van der Werf et al., 2017; van der Werf et al., 2010). Uncertainties in GFED4.1s stem from uncertainties in BA, (see above), the amount of biomass consumed per unit BA, and the carbon emitted per unit biomass burned. Revisions to BA input data, discussed above, have tended to influence GFED central estimates of fire C emissions to a greater degree than the uncertainties around central estimates (van der Werf et al., 2017; Chen et al., 2023). Uncertainties in fire carbon emissions estimates from GFAS are on the order of approximately $\pm 25\%$ at 1 standard deviation for global totals. Uncertainties are introduced by missed active fire detections, either below the detection threshold of the MODIS instruments, or not observed during the limited diurnal coverage of Low Earth Orbiting satellites, assumptions made for biome classifications, coefficients used to convert observed thermal anomalies to consumed dry matter, and emission factors used to estimate emitted quantities of carbon and pyrogenic pollutants. Variation in C emissions estimates on the order of approximately 20-60% has been observed in studies comparing multiple emissions products (Wiedinmyer et al., 2023).

The fire radiative power (FRP) data provided by the MOD14A1 and MYD14A1 products are subject to several well-documented qualitatively-described uncertainties that affect both the detection of active fires and the precision of retrieved energy estimates (Giglio et al., 2016; Wooster et al., 2021). Omission errors typically arise when fires are obscured by clouds or, in

some cases, dense smoke incorrectly flagged as clouds during masking procedures (Atwood et al., 2016). Additional omissions occur when the mid-infrared (MIR) radiance levels of small, low-intensity fires fall below detection thresholds, which is most common in the case of sub-canopy or peatland combustion (Schroeder et al., 2008; Roberts et al., 2018). Temporal gaps in satellite coverage also contribute, as MODIS instruments observe any given location only up to four times per day, often missing short-lived events or peak fire activity in the late afternoon (Roberts and Wooster, 2014). Commission errors, by contrast, typically occur when non-fire thermal anomalies are misclassified as active fires. False positives can be caused by sunglint on water or clouds or by thermally anomalous surfaces such as bare soils, urban infrastructure, gas flares, and volcanic eruptions, which produce elevated MIR radiance that mimics fire signatures (Wooster et al., 2021). Contextual detection algorithms help mitigate these errors by comparing candidate pixels to local background conditions. These approaches have been particularly successful in reducing commission errors, which are often below 10% (Giglio et al., 2016; Wooster et al., 2021). In contrast, uncertainties in omission errors and FRP observations remain less well characterised (Wooster et al., 2021; Li et al., 2024). [Freeborn et al. \(2013\) quantified overall MODIS FRP uncertainty by analysing >400,000 near-simultaneous duplicate fire detections from consecutive satellite scans, comparing overlapping measurements to assess variability mainly in relation to a fire's position within the sensor's point spread function. They found uncertainties of ~27% \(at 1 standard deviation\) for individual fire pixels, but showed through simulations that these decline rapidly with aggregation, falling to ~17% for two-pixel clusters and below 5% once ~50 or more pixels are included, making large-fire or regional FRP estimates more robust than single-pixel estimates.](#)

2.1.1.3. *Regional Burned Area, Carbon Emissions and Fire Count Totals*

We calculated regional totals of BA and C emissions based on a variety of regional layers defined in **Table 1**. The regional layers represent a range of biogeographical boundaries (e.g. biomes), geopolitical boundaries (e.g. countries), and values used in scientific reports (e.g. by the Intergovernmental Panel on Climate Change; IPCC). We calculated monthly totals of BA and fire C emissions for each region by aggregating monthly BA and daily C emissions data, summing the data from the input datasets both spatially and temporally as required. In the case of fire C emissions, we also calculated the mean estimate of fire C emissions from GFED4.1s and GFAS, regionally.

~~We adopt a~~Our March-February definition of the global fire season (e.g. the latest global fire season spans March 2024-February 2025). ~~Due to~~ is chosen so as to align with an annual lull in the global fire calendar in the boreal spring months (Giglio et al., 2013). ~~It was previously shown that~~ fire season BA totals are least sensitive to the shifts in fire season cutoffs of 1-2 months if the fire season centres on boreal spring (Boschetti and Roy, 2008). This makes the global fire season centred on spring a pragmatic option for the study of interannual variability or trends in fire extent (Boschetti and Roy, 2008). The period March-February is specifically oriented at the end of the austral fire season and before widespread fires have begun in the boreal extratropics. The regions where this global definition of the fire season is most problematic are: northern hemisphere South America, Southeast Asia, and Central America (Giglio et al., 2013).

In addition, we calculated totals of regional fire counts for each global fire season based on the number of individual fire ignition points present within each region, using ignition point vectors from the Global Fire Atlas (Andela et al., 2019). The resolution of the MODIS MCD64A1 data supplied to the Global Fire Atlas algorithm is 500 m and hence fires ~~that are smaller in scale are~~ small fires omitted from the MCD64A1 product are also omitted from the Global Fire Atlas outputs. Regional or national systems may record greater fire counts due to the inclusion of smaller fires.

Table 1: Regional layers to which global Earth observations were disaggregated and used to define regions with extreme wildfire seasons or extreme individual wildfire attributes. Regional layers are available from Jones et al. (2025).

Layer	Short Form	Source	Notes
Biomes	NA	Olson et al. (2001)	
Continents	NA	ArcGIS Hub (2024)	
Continental Biomes	NA	Olson et al. (2001), ArcGIS Hub (2024)	Spatial intersect of biomes and continents.
Ecoregions	NA	Olson et al. (2001)	Ecoregions are geographically inset within biomes.
Countries	NA	EU Eurostat (2020)	
UC Davis Global Administrative Areas (GADM) Level 1	GADM-L1	UC Davis (2022)	First sub-national administrative level, such as states of the US or provinces of China. Version 4.1.
Intergovernmental Panel on Climate Change <i>Sixth Assessment Report (AR6) Working Group I (WGI) Reference Regions</i>	IPCC AR6 WGI Regions	Iturbide et al. (2020)	
Global C Project <i>Regional C Cycle Assessment and Processes (RECCAP2) Reference Regions</i>	RECCAP2 Regions	Ciais et al. (2022)	
Global Fire Emissions Database (GFED) Basis Regions	GFED4.1s Regions	van der Werf et al. (2006)	

2.1.1.4. Cross-Product Intercomparison of Regional Burned Area Totals

In this report, to characterise the dependence of our findings on BA product choices, we add a supplementary comparison between the regional BA totals detected by the MCD64A1 BA product and two other BA products. The first product was the ESA Climate Change Initiative FireCCIS311 product, derived from Sentinel-3 SYN reflectance and Visible Infrared Imaging Radiometer Suite (VIIRS) active fires (Lizundia-Loiola et al., 2022; see **Figure S1**). FireCCIS311 is provided at a spatial resolution of 300 m and is based on a contextual algorithm based on Sentinel-3 SYN surface reflectance (SYN combines OLCI and SLSTR reflectance), guided by active fire detections from VIIRS. The second product is NASA’s VIIRS BA product (VNP64A1 v002; Zubkova et al., 2024; Giglio et al., 2024; see **Figure S1**), generated using an adaptation of the MODIS MCD64A1 Collection 6.1 algorithm, applied to 750 m VIIRS imagery and active fire detections. The hybrid algorithm uses dynamic thresholds on composite imagery derived from a burn-sensitive vegetation index and temporal texture measures, enabling it to distinguish fire-induced changes from other land surface changes. It identifies the burn date at 500 m resolution for each grid cell, with prior probabilities of burned/unburned areas informed by cumulative VIIRS active fire observations.

The FireCCIS311 product has been computed since 2019, and hence our cross-product comparisons focus on the fire seasons March 2019-February 2025. We followed identical

approaches as described in prior sections to calculate regional BA totals and to quantify anomalies of the past fire season. With very few exceptions, we find a high level of consistency between the MCD64A1, FireCCIS311, and VIIRS VNP64A1 BA products with regards to both the relative magnitude of regional BA ~~totals and anomalies~~, the geographical distribution of those anomalies, and the rankings of BA in the 2024-25 fire season versus previous fire seasons since 2019 (**Figure S1**; Jones et al., 2025). This analysis adds confidence that regional anomalies identified in the MCD64A1 BA product are generally replicated consistent across products from different space agencies using different algorithms applied to different combinations of Earth-observing sensors. The MCD64A1 BA product will soon discontinue due to the decommissioning of MODIS sensors aboard NASA's Terra and Aqua satellites. Consistency, and The consistency across products is an encouraging finding for the continuity of our annual reporting.

2.1.2. Identifying Extreme Fire Seasons and Events from Earth Observations

2.1.2.1. Regions with Extreme Wildfire Seasons

Anomalies in BA, fire C emissions, and fire counts in the latest global fire season (March 2024-February 2025) were calculated in several ways:

- (i) as relative anomalies (expressed in %) from the annual mean during all previous March-February periods since 2002 (2003 for fire C emissions);
- (ii) as standardised anomalies (standard deviations) from the annual mean during all previous March-February periods since 2002 (2003 for C emissions);
- (iii) as a rank amongst all March-February periods since 2002 (2003 for fire C emissions), March 2024-February 2025 inclusive.

In this report, anomalies in fire C emissions are reported based on the two-model mean estimate from GFED4.1s and GFAS, however anomalies based on the GFED4.1s or GFAS estimates individually are also available via Jones et al. (2025).

We identified regions in which the latest fire season was potentially classifiable as 'extreme' based on the rank of BA, C emissions, and fire count amongst all fire seasons. For visualisation purposes, we identified regions in which the latest fire season ranked in the top 5 of all annual fire seasons on record (see **Section 2.2.1**). The BA data for the period March 2002-February 2025 includes 23 fire seasons, while the C emissions data for the period March 2003-February 2025 includes 21 fire seasons. Hence, a top-5 ranking translates approximately to a fire season in the upper quartile of those on record.

We further characterised the onset, peak, and cessation of anomalous monthly BA in March 2024-February 2025. First, we identified the 'peak' month ~~of the event's peak~~ as the maximum difference anomaly between monthly BA values in March 2024-February 2025 and the climatological mean monthly values from the prior March-February periods. Thereafter, the event's onset and cessation were defined as the bounds of consecutive months with above-average BA prior to and following the peak but limited to the March 2024-February 2025 period.

The annual data and anomalies produced using these methods are available from Jones et al. (2025).

2.1.2.2. Regions with Extreme Individual Wildfire Attributes

We identified regions in which large or fast-moving fires occurred in the latest fire season based on records of individual fires from the Global Fire Atlas (Andela et al., 2019). For each

region (**Table 1**) and year, we estimated the size of the largest fire, the daily rate of growth of the fire that spread most rapidly, the size of the 95th percentile fire, and the daily rate of growth of the 95th percentile fire. In the Global Fire Atlas, the daily rate of growth for any given fire is determined by calculating the average daily rate of growth at which the fire advanced across all its constituent cells. ~~This method includes cells burned by the head, flank, and backfire and produces lower spread rates than if the calculation were based solely on the cells burned by the head fire.~~

As a new analysis developed for the 2024-25 report, we also identified regions in which intense fires occurred in the latest fire season based on the Global Fire Atlas and FRP observations from the MODIS active fire datasets (MOD14A1 and MYD14A1). Regional values were calculated per fire season across two steps as follows. First, each fire present in the Global Fire Atlas was assigned a peak intensity value equivalent to the 95th percentile of all FRP measurements (daytime and nighttime) occurring within the perimeter and date range of the fire. Second, the regional summary values were taken to be the mean of all peak (95th percentile) intensity values from the cohort of fires occurring in a region and fire season. This approach effectively masks FRP measurements to fires that occur in the Global Fire Atlas prior to averaging, meaning that the fire intensity anomalies presented here relate to the same set of fires as the fire size and fire rate of growth statistics.

Anomalies in each fire attribute were calculated relative to other fire seasons since 2003 using the same metrics as for BA (see *i-iii* above), and we identified regions in which the latest fire season featured fires with potentially extreme attributes based on the ranking of the individual fire metrics amongst all fire seasons.

The annual data and anomalies produced using these methods are available from Jones et al. (2025).

2.1.3. Identifying Extreme Fire Seasons and Events from Expert Consultation

2.1.3.1. Role of Expert Consultation

We assembled a panel of regional experts from each continent (**Table A1**) to contribute to the identification, description, and characterisation of extreme wildfire seasons or impactful events in the latest fire season. A key role of the expert panel was to catalogue regional events that significantly impacted society or the environment but which may not have been detected by Earth-observing satellites due to issues such as scale, short duration, timing of overpass, and cloud or canopy cover. This includes (but is not limited to) wildfires that presented major suppression difficulty, impacted society by causing fatalities, evacuations, displacement (e.g. homelessness), direct structure or infrastructure loss or damage, degradation of air or water quality, loss of livelihood, cultural practice or other ways of life, and loss of economic productivity. This definition also includes (but is not limited to) wildfires that impact the environment via disturbance to vulnerable ecosystems, biodiverse areas, or ecosystem services such as C storage. This approach recognises that Earth observations do not provide a complete record of all impactful fires. We do not define ubiquitous quantitative thresholds of impact by any of the measures outlined above, but rather invite in-region experts to identify events that triggered impacts that were sufficient in magnitude to infiltrate public and political discourse. The sources of information available for cataloguing regional events include national/regional fire records, land and fire management agencies reports, disaster management reports, news reports, and social media. A second key role of our expert panel was to describe and contextualise the impacts of the fire seasons highlighted as extreme by Earth observations or regional assessment (see **Section 2.2.3**). We stress that extremes identified in this way often reflect vulnerability and resilience shaped by factors beyond climate.

such as governance, funding, and policy decisions, which are not strictly biophysical conditions.

The year in review by continent, produced by the expert panel, is presented in **Appendix A**.

2.1.3.2. Shortlisting of Focal Events

In later sections of this report, we conducted various analyses to understand the causes and predictability of a selection of extreme wildfire seasons or events during March 2024-February 2025 (see **Sections 4-6**). We limited the number of analyses to three globally prominent focal events of the 2024-25 global fire season because the approaches used are not operational and time is required to train and optimise our models regionally.

In discussion with our expert panel, we prioritised the three events studied in this report by weighing up the anomalies in Earth observations during the latest fire season as well as a suite of impacts that these extremes had on people and the environment. The focal events are notable for their international significance even where they have not attracted international media attention and where they have been highly relevant and recognized within and beyond their region.

2.1.4. Contextualising Analyses

2.1.4.1. Contemporaneous Extremes in Fire Weather

In the supplementary material edition of this report, we introduce routine summaries of the extreme (95th percentile) fire weather days during the March 2024-February 2025 global fire season based on [the Fire Weather Index \(FWI\)](#), the FWI, a common metric of fire danger developed by the Canadian Forest Service as part of the Canadian Forest Fire Danger Rating System (CFFDRS; van Wagner, 1987). The FWI comprises various components that consider the influence of weather on fire danger, with 2m temperature, 10m wind speed, precipitation, and 2m relative humidity as prerequisite variables. Higher FWI values are generally seen during droughts, heatwaves and strong winds as these conditions are conducive to wildfires in environments with sufficient fuel load (Jolly et al., 2015; Di Giuseppe, 2016; Jones et al., 2022). We base our analysis of extreme (95th percentile) fire weather on the FWI dataset derived from the Copernicus Climate Change Service ERA5 reanalysis (Hersbach et al., 2020; Vitolo et al., 2020) and maintained by the Copernicus Emergency Management Service (CEMS, version 4.1, 2019). [The 95th percentile value for extreme fire weather has been used in many prior publications \(e.g. Abatzoglou et al., 2019; Jones et al., 2022\)](#). The same statistics are reported for the 2024-25 fire season as in the case of fire observational datasets, including (i) ranks, (ii) proportional anomalies, and (iii) standardised anomalies amongst all fire seasons since 2002 (**Figure S2**). Full discussion of the methodology and results are provided in **Supplementary Text S2**. The data produced using these methods are available from Turco et al. (2025).

2.1.4.2. 21st Century Trends in Burned Area

To place recent extremes in the context of fire trends of the past two decades, we update our regional analyses of trends in annual BA from Jones et al. (2022). In addition to reporting trends in *total* BA, we also present trends in *forest* BA as these regularly diverge from total BA trends (**Figure S3**), following Jones et al. (2024a). Full discussion of the methodology and results are provided in **Supplementary Material S2**.

2.2. Results

2.2.1. Extreme Fire Seasons and Events of 2024-25 from Global Earth Observations

2.2.1.1. Global Summary

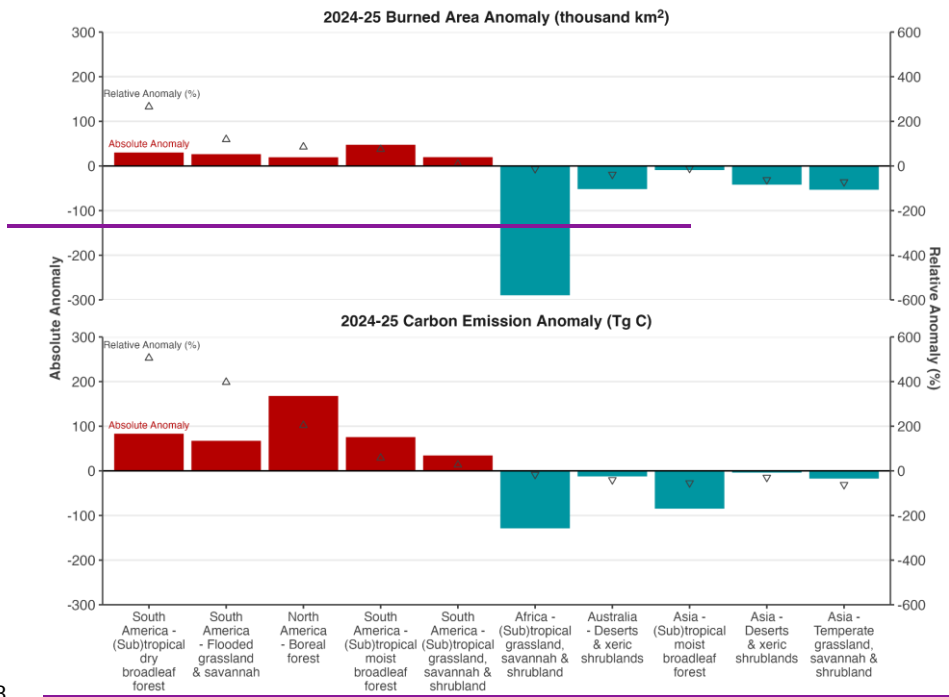
According to the MODIS BA product, at least 3.7 million km² burned globally during the 2024-25 global fire season (March 2024-February 2025), 9% below the average of previous fire seasons (4.0 million km²) since 2002 and overall ranking 16th (i.e., 8th lowest) of all fire seasons since 2002 (Jones et al., 2025). Despite this, fire C emissions were 9% above average at 2.2 Pg C during the 2024-25 global fire season, which ranks 6th amongst all fire seasons since 2003 (based on annual averages of GFED4.1s and GFAS estimates; see **Section 2.1.2**; Jones et al., 2025). The 2024-25 fire season therefore followed a similar pattern as in the 2023-24 fire season, with above-average emissions occurring despite below-average BA at the global level. These anomalies, signifying lesser fire extent but more severe fires than average, are emblematic of consistent with a reported trend towards increased fire extent and intensity in carbon-rich environments such as forests globally (Jones et al., 2024a). It is important to note that the MODIS BA product is uncorrected for missed small fire detections as in the case other estimates (e.g. Chen et al., 2023; Lizundia-Loiola et al., 2022), meaning that the estimated BA extents from MODIS are conservative (i.e., at least 3.7 million km² burned globally during 2024-25).

Stark regional contrasts in the anomalies in BA, fire C emissions and individual fire properties are visible in the Earth observations at various regional scales (**Figure 1, Figure 2, Figure 3**). The three countries with greatest positive anomalies in BA and C emissions during 2024-25 were Bolivia, Brazil, and Canada (**Table 2, Table 3**), marking a second consecutive year in which the Americas experienced an anomalous fire season.

On the scale of continental biomes (**Figure 1, Figure 2, Figure 3**), the greatest BA and fire C emissions anomalies of 2024-25 were seen in the North American boreal forests (mostly in Canada), the South American moist tropical forests (mostly in Amazonia), the South American dry tropical forests (mostly in the Chiquitano dry forests of Bolivia), and the South American grassland and savannah biome (mostly in the Cerrado region). On the other hand, it was a second consecutive year the African savannahs experienced a low fire season. In the world's tropical savannah regions, which contribute around 70% towards global BA, the total BA in the 2024-25 fire season was 290 thousand km² (12%) below average in Africa, slightly above average in South America, and slightly above average in Australia (**Figure 2**). Total BA across the global (sub)tropical grassland, savannah, and shrubland biome was 290 thousand km² (10%) below average, and the 6th lowest on record, but still contributed 70% towards total global BA during 2024-25. Correspondingly, the C emitted by fires in global savannahs was 102 Tg C (10%) below average in 2024-25.

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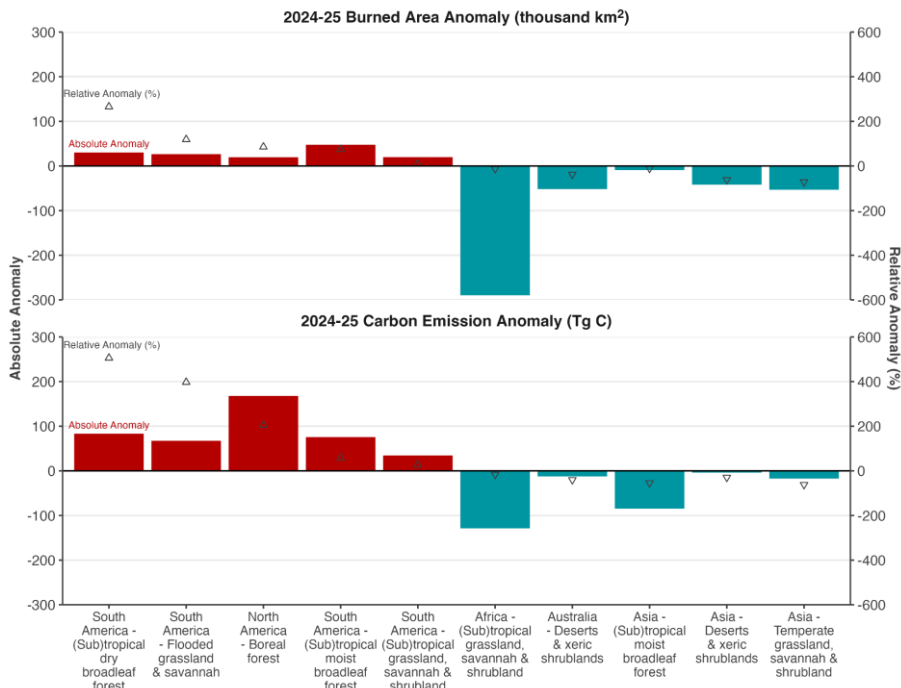


Figure 1: Anomalies in burned area (BA) and carbon (C) emissions for selected continental biomes in the 2024-25 global fire season (March 2024-February 2025), versus the average of prior fire seasons since 2002. The selected regions all experienced BA anomalies of over ± 20 thousand km² or C emissions anomalies over ± 30 Tg C during the 2024-25 global fire season. Relative changes (%) are also marked by triangular symbols and can be read on the secondary axis. The notation "(sub)tropical" is an abbreviation of "tropical and subtropical" and is used consistently in this report.

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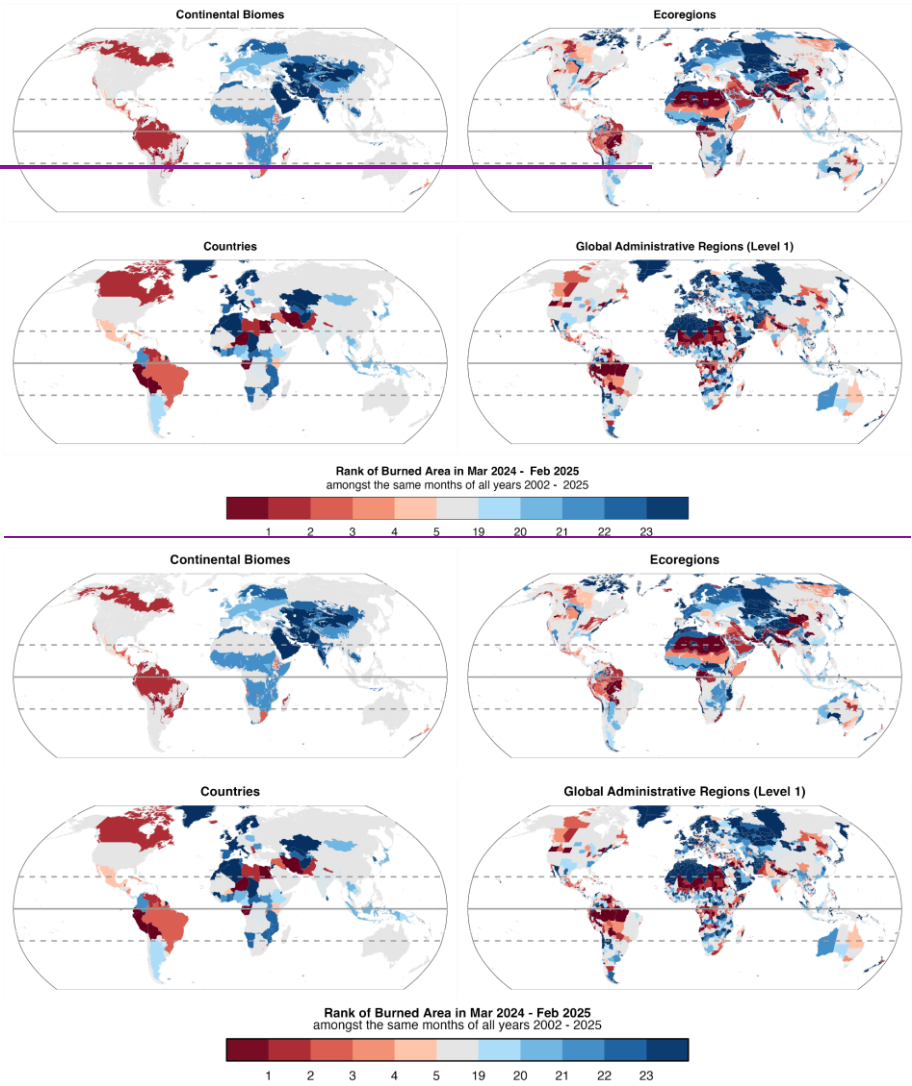


Figure 2: Ranks of BA during March 2024-February 2025 versus previous March-February periods (n = 23 global fire seasons), at the scales of (top left) continental biomes, (top right) ecoregions, (bottom left) countries and (bottom right) level 1 administrative regions. Results for regions with high-ranking (top 5 years) or low-ranking (bottom 5 years) events are highlighted. The timing of BA anomalies is shown in **Figure S4**.

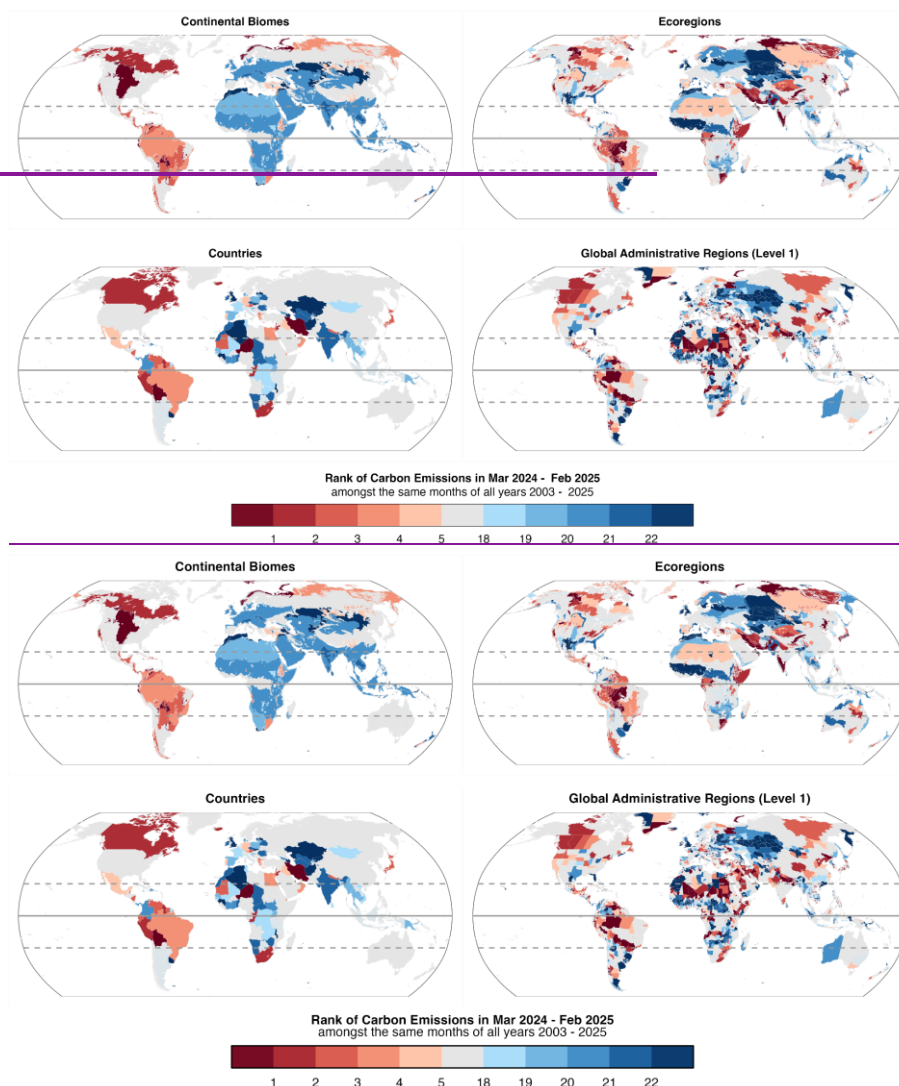


Figure 3: Rank of fire C emissions during March 2024-February 2025 versus all March-January periods since 2003 (n = 22 global fire seasons), at the scales of (top left) continental biomes, (top right) ecoregions, (bottom left) countries and (bottom right) level 1 administrative regions. We consider C emissions estimates from two products (GFAS and GFED), first calculating the mean emissions value from the two products, then ranking the values.

Table 2: Summary of the largest positive anomalies in burned area (BA) during the 2024-25 fire season on national and sub-national scales. Anomalies are expressed relative to all previous fire seasons 2002-2024 (n = 23). The table includes the top ten countries ranked by the magnitude of their absolute BA anomalies and the top 30 level 1 administrative regions

968 (e.g. states or provinces) grouped into countries where applicable. Extended data for all
969 countries and region layers are available from Jones et al. (2025).

Region Name	BA during the 2024-25 fire season (thousand km²)	Absolute BA anomaly (thousand km²)	Relative BA Anomaly (%)	Ranking of the 2024-25 fire season
Bolivia	107	+67	+169	1
Santa Cruz (Department of Bolivia)	65	+49	+311	1
Beni (Department of Bolivia)	36	+15	+74	4
Brazil	243	+59	+32	3
Mato Grosso (State of Brazil)	68	+22	+49	4
Pará (State of Brazil)	36	+20	+119	1
Mato Grosso do Sul (State of Brazil)	23	+11	+90	2
Amazonas (State of Brazil)	9	+6	+254	1
São Paulo (State of Brazil)	10	+4	+67	4
Canada	46	+21	+86	2
Northwest Territories (Territory of Canada)	16	+12	+281	3
British Columbia (Province of Canada)	8	+5	+154	4
Alberta (Province of Canada)	7	+4	+123	2
Venezuela	43	+15	+52	2
Apure (State of Venezuela)	16	+5	+41	2
Bolívar (State of Venezuela)	6	+3	+133	1
Niger	13	+10	+257	1
Tahoua (Department of Niger)	5	+4	+967	1
Burkina Faso	33	+9	+39	5
Sahel (Region of Burkina Faso)	6	+6	+1226	1
Angola	374	+9	+2	8
Moxico (Province of Angola)	61	+8	+15	3
Huíla (Province of Angola)	20	+6	+49	1
Cunene (Province of Angola)	18	+5	+35	5
Bié (Province of Angola)	20	+4	+25	1
Congo (Republic of the)	41	+8	+25	1
Sudan	82	+8	+11	8
North Darfur (State of Sudan)	15	+9	+168	1
Mali	77	+7	+10	6
Gao (Region of Mali)	13	+12	+1383	1
Other				
Queensland (State of Australia)	100	+19	+24	5
Heilongjiang (Province of China)	23	+14	+164	2
Zabaykalye (Territory of Russia)	23	+11	+88	3
North-Western (Province of Zambia)	45	+10	+29	1
Sakha (Republic of Russia)	27	+9	+55	6
Amur (Region of Russia)	20	+8	+70	4
Zamfara (State of Nigeria)	9	+5	+95	4
Oregon (State of United States)	7	+5	+285	1
Jilin (Province of China)	7	+4	+186	4
Sankuru (Province of Dem. Rep. Congo)	11	+4	+58	1

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Table 3: Summary of the largest positive anomalies in carbon (C) emissions during the 2024-25 fire season on national and sub-national scales. Anomalies are expressed relative to all previous fire seasons 2003-2024 (n = 22). The table includes the top ten countries ranked by the magnitude of their absolute C emissions anomalies and the top 30 level 1 administrative regions (e.g. states or provinces) grouped into countries where applicable. Extended data for all countries and region layers are available from Jones et al. (2025).

Region Name	C emitted during the 2024-25 fire season (Tg C)	Absolute C emissions anomaly (Tg C)	Relative C emissions Anomaly (%)	Ranking of the 2024-25 fire season
Canada	282	+189	+204	2
Northwest Territories (Territory of Canada)	104	+85	+441	2
Alberta (Province of Canada)	56	+42	+297	2
British Columbia (Province of Canada)	55	+36	+196	2
Saskatchewan (Province of Canada)	43	+28	+184	3
Manitoba (Province of Canada)	11	+5	+74	4
Bolivia	187	+148	+383	1
Santa Cruz (Department of Bolivia)	157	+136	+637	1
Beni (Department of Bolivia)	23	+11	+86	3
La Paz (Department of Bolivia)	4	+2	+79	4
Brazil	314	+111	+55	4
Mato Grosso (State of Brazil)	86	+29	+50	6
Amazonas (State of Brazil)	35	+25	+237	1
Mato Grosso do Sul (State of Brazil)	30	+23	+323	1
Par�� (State of Brazil)	59	+22	+61	4
Tocantins (State of Brazil)	22	+5	+33	5
S��o Paulo (State of Brazil)	8	+5	+190	1
Rond��nia (State of Brazil)	22	+3	+16	7
Roraima (State of Brazil)	5	+2	+81	5
Venezuela	26	+8	+47	3
Bol��var (State of Venezuela)	5	+2	+97	1
Mexico	29	+6	+26	5
South Africa	18	+3	+24	2
Angola	146	+3	+2	9
Moxico (Province of Angola)	28	+5	+21	3
Bi�� (Province of Angola)	9	+2	+35	1
Hu��la (Province of Angola)	7	+2	+37	1
Peru	7	+2	+51	2
Russian Federation	179	+2	+1	9
Sakha (Republic of Russia)	75	+32	+74	3
Zabaykal'ye (Territory of Russia)	31	+14	+78	4
Amur (Region of Russia)	25	+8	+46	5
Arkhangel'sk (Region of Russia)	2	+2	+1776	1
Congo (Republic of the)	10	+2	+24	2
Other				
Queensland (State of Australia)	31	+4	+14	7
Oregon (State of United States)	7	+4	+130	3
Idaho (State of United States)	5	+3	+139	3
North-Western (Province of Zambia)	22	+2	+12	1
Alto Paraguay (Department of Paraguay)	6	+2	+55	2
Mai-Ndombe (Province of Dem. Rep. Congo)	7	+2	+36	1

2.2.1.2. An Unprecedented Fire Season in South America

There were pronounced and widespread positive anomalies in BA in 2024-25 across South America during 2024-25 (**Figure 1, Figure 2**). Several South American biomes experienced extremely high or even record-setting BA in the 2024-25 fire season (**Figure 1**). The South American (sub)tropical dry broadleaf forests, principally comprising the Chiquitano and Chaco dry forests, experienced a record-breaking fire season, with the 42 thousand km² burned exceeding the average since 2002 by a factor of 3.6 and the 100 Tg C emitted exceeding the average since 2003 by a factor of 6. In the South American (sub)tropical moist broadleaf forests, principally comprising the Amazon rainforest, BA was 47 thousand km² (75%) above the average since 2002, which is the second-highest year on record, and C emissions were correspondingly 76 Tg C (58%) above average. Finally, in the South American Flooded grassland and savannah biome, which principally includes the seasonally inundated Pantanal region, BA was 26 thousand km² (119%) above the average since 2002, which is also the second-highest year on record, and C emissions were correspondingly 67 Tg C (397%) above average. Across South America as a whole, BA was 120 thousand km² (35%) above average and C emissions were 263 Tg C (84%) above average, producing the highest C emissions total on record for the continent.

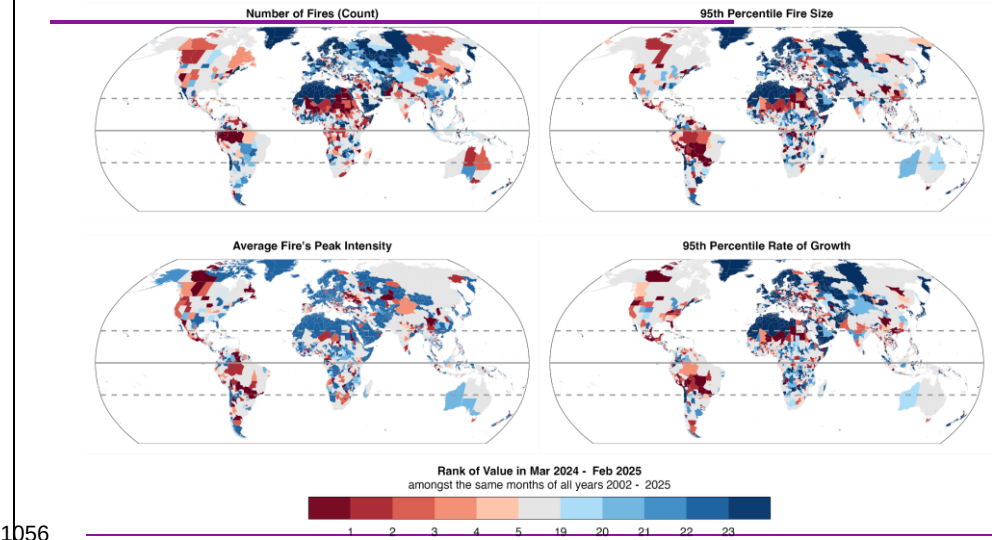
The spatial breadth of the record-setting or high-ranking anomalies in fire extent, emissions, size, rate or spread and intensity (**Figure 2, Figure 3, Figure 4**), as well as their impact on society and the environment, made the last fire season unprecedented on the continent. **Appendix A (Section A6)** discusses the unprecedented South American fire season of 2024-25 in greater detail, including its impacts and regional context, relying also on information from regional fire monitoring systems and reporting.

Fifteen of South America's 115 ecoregions experienced new record levels of BA or C emissions during 2024-25 (**Figure 2, Figure 3**) and 72 of South America's ecoregions experienced BA or C emissions in the top three years on record (**Figure 2, Figure 3**). Regions with record levels of BA or C emissions included the Chiquitano dry forests and the Pantanal wetlands of Bolivia and central-west Brazil. In nearby southern and southwestern parts of Amazonia, five moist forest and seasonally flooded (várzea) ecoregions also showed record-breaking BA or C emissions. The widespread positive BA anomalies in southern and southwest Amazonia, the Chiquitano and the Pantanal were visible in the MODIS BA dataset from March and April 2024, peaking in August-November 2024 before subsiding around November (**Figure S4**). In the Guianan shield region, encompassing much of Northeast Amazonia (north of the Amazon river and the Rio Negro tributary) and the Guianan forests of Venezuela, Guyana, and Suriname, four moist forest and swamp forest ecoregions also experienced record-breaking levels of BA or C emissions (**Figure 2, Figure 3**). Here, BA anomalies peaked around March-April before subsiding in May in northern parts but persisted through to December in areas closer to the equator (**Figure S4**).

At the national level within South America, the most significant anomalies in BA during the 2024-25 fire season occurred in Bolivia, where BA was 67 thousand km² (169%) above average and fire C emissions were 148 Tg C (383%) above average, the greatest values on record in the country (**Figure 2, Figure 3; Table 2, Table 3**). In Brazil, BA was 59 thousand km² (32%) above average and emissions were 111 Tg C (55%) above average during 2024-25, making it the country's third highest fire season on record for BA after 2007-08 and 2010-11. Additionally, Venezuela recorded an anomaly of +15 thousand km² (+52%), its second-highest BA total after 2023-24. Anomalies in these three countries are highlighted due to global totals of BA and C emissions (**Table 2, Table 3**). On sub-national scales, the 2024-25 fire season saw record-breaking BA or C emissions in four states of Brazil (Pará, Amazonas, Mato Grosso do Sul, and São Paulo), one department of Bolivia (Santa Cruz), 3 States of Venezuela (Bolívar, Delta Amacuro, Monagas). Other record-breaking anomalies were seen at sub-national levels across South America (**Figure 2, Figure 3**), including in 6 regions of Guyana,

1033 7 regions of Peru, 2 districts of Suriname, 8 provinces of Ecuador, as well as some parts of
1034 Chile and Colombia (**Figure 2, Figure 3**), clearly signalling the large geographical breadth of
1035 the extremes on the continent during the 2024-25 fire season.
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1037 For most regions of South America, the anomalies in BA and C emissions were explained by
1038 particularly large, fast moving and intense fires, rather than above-average fire counts (**Figure**
1039 **4**). In Brazil, data on individual fire characteristics from the Global Fire Atlas showed new
1040 record fire sizes at the 95th percentile threshold for 6 states (Amapá, Mato Grosso, Mato
1041 Grosso do Sul, Paraná, Rondônia, and São Paulo). In Mato Grosso, Mato Grosso do Sul, and
1042 São Paulo, 95th percentile fire sizes were 105-266% above average, driving record breaking
1043 BA despite fire counts being 18-54% below average. Meanwhile, three states (Mato Grosso,
1044 Mato Grosso do Sul, and São Paulo) all saw the fastest rates of growth at the 95th percentile
1045 threshold, and 5 states (Mato Grosso do Sul, Paraná, Rio de Janeiro, Roraima, and São
1046 Paulo) experienced the most intense fires on record (measured per the average fire's 95th
1047 percentile intensity-value; **Figure 4**). Unlike in other parts of Brazil, the fire count anomaly
1048 (+154%) was record-breaking in Amazonas during 2024-25, combining with the 95th
1049 percentile fire size anomaly (+60%) to produce the record-breaking BA. Similar patterns were
1050 observed across South America, with anomalies in fire size, rates of growth, and intensities
1051 generally being more widespread than anomalies in fire count (**Figure 4**). Some notable
1052 exceptions were 5 regions of Peru, 5 regions of Ecuador, 3 regions of Colombia, and 3 regions
1053 of Guyana, where record-setting fire counts were observed, as well as in parts of Venezuela
1054 where high-ranking fire counts occurred (**Figure 4**).
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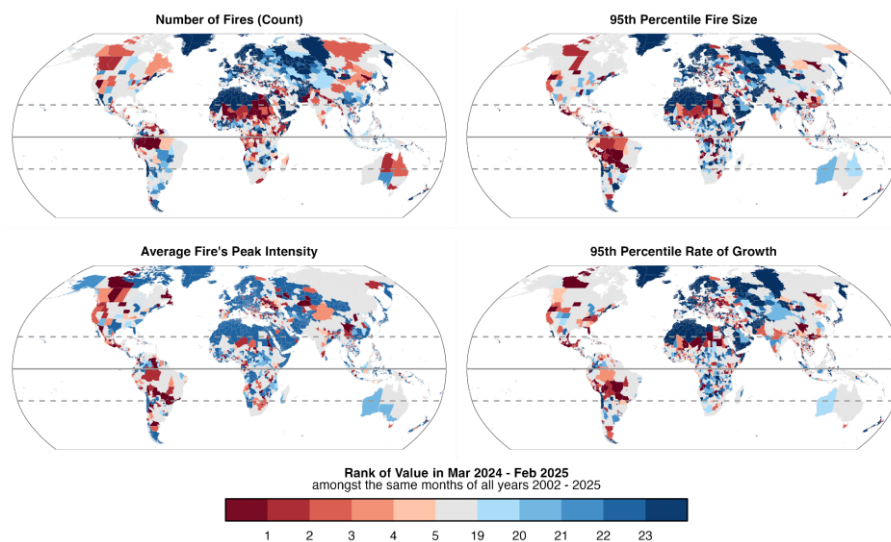


Figure 4: Ranks of selected individual fire properties during the March 2024-February 2025 fire season versus previous March-February periods ($n = 23$ global fire seasons), including **(top left)** fire count, **(top right)** 95th percentile fire size, **(bottom left)** the average value of a the peak intensity (95th percentile FRP within fire perimeters) considering all regional fires, and **(bottom right)** 95th percentile daily rate of growth. Results are shown at the scale of states or provinces (GADM administrative level 1 regions).

2.2.1.3. A Second Consecutive Extreme Fire Year in North America

The 2024-25 fire season was the second-highest fire year on record for BA and C emissions in the North American boreal forests, with BA 86% above the average since 2002 (+20 thousand km²) and C emissions 3 times the average since 2003 (+168 Tg C). These large anomalies follow the record-breaking 2023-24 fire season when BA was five times above average and C emissions were seven times above average, marking two consecutive years of extreme fire activity in the North American boreal forests. Elsewhere, BA extent was in the top three years on record in the North American (sub)tropical moist broadleaf forest (concentrated in Latin America), and in the North American mediterranean forests, woodlands and scrub (concentrated in Southern California). Across North America as a whole, BA was 31 thousand km² (35%) above average and C emissions were 194 Tg C (112%) above average, the second highest totals on record for both metrics.

Eleven of North America's 189 ecoregions experienced new record levels of BA or C emissions during 2024-25 (**Figure 2, Figure 3**), with these regions principally concentrated in northwest Canadian taiga and tundra, mountain forests of the northwest US and southwest Canada (principally in Oregon and Alberta), and moist tropical forest ecoregions of mesoamerica (principally in Mexico), but also including the Central Valley grasslands of California and the northeast coastal forests of the US. More broadly, but with a similar geographical distribution, 44 North American ecoregions experienced BA or C emissions in the top three years on record (**Figure 2, Figure 3**). The positive BA anomalies in extratropical North America were visible in the MODIS BA dataset from April 2024 in western regions (e.g. mountain forests of the northwest US and southwest Canada), July-August 2024 in the central regions (e.g. Canadian tundra and taiga), and late into the 2024 summer in eastern regions (e.g. northeast coastal forests; **Figure S4**). Thereafter, BA anomalies were consistently

observed through summer (July-September 2024) and in some cases persisted through October 2024.

In Canada, BA was 21,000 thousand km² (86%) above average and C emissions were 189 Tg C (204%) above average during 2024-25, marking the country's second highest fire season on record immediately following the record-breaking fire season of 2023-24 (**Figure 2, Figure 3; Table 2, Table 3**). Notably, the anomalies of 2024-25 were concentrated in the western Canadian states of British Columbia, Alberta and northwest Territories which all saw the second-highest BA or C emissions on record, with large anomalies in the range of 120-440%, second only to the 2023-24 fire season. More generally, record levels of BA or C emissions were less spatially extensive in North America than in South America, though the US states of Oregon, Wyoming, and New York saw record BA, as did several mesoamerican states of Mexico, Guatemala, and Costa Rica (**Figure 2, Figure 3**).

For western Canada, individual fire metrics from the Global Fire Atlas were also anomalous and highly-ranked amongst previous years, but generally fell short of the records set in the 2023-24 fire season (**Figure 4**). For example, fire counts were 170-190% above average in Alberta and British Columbia, ranked second (behind 2023-24), whereas anomalies in 95th percentile fire size and rate of growth were not particularly large. Meanwhile, the explanation for the anomalous BA in some states of the northwest US was not consistent, with some states experiencing above-average fire counts, some experiencing above-average fire sizes, but few experiencing both.

Appendix A (Section A4) provides a more complete summary of the fire season in North America based on the regional panel assessment.

2.2.1.4. A Mixed Picture in Africa

For the second consecutive year, BA was around 290 thousand km² (12%) below the average of previous fire seasons in the African (sub)tropical grassland, savannah and shrubland biome, and the 3rd lowest on record (**Figure 2**), but still contributed 56% towards the global BA total and 86% towards total BA in Africa. BA anomalies in the African savannahs have a significant influence on the continental BA anomalies, and indeed BA across Africa as a whole was 313 thousand km² (12%) below average.

Despite the low fire activity in Africa during 2023, several exceptions emerged in both central and northern Africa. Record levels of BA were observed in several parts of the Congo Basin (**Figure 2, Figure 3**) due to an unusually high number of fires (**Figure 4**). BA in the Republic of Congo was 25% above average, the highest on record, and similarly fire C emissions were 25% above average (**Table 2, Table 3**). In the Democratic Republic of the Congo, the Mai-Ndombe and Sankuru provinces each experienced record levels of BA or fire C emissions with anomalies in the range of 36-58% (**Table 2, Table 3**). These anomalies were centred on several western ecoregions of the Congo Basin, including the Atlantic Equatorial coastal forests where BA was more than triple the annual mean, Western Congolian swamp forests where BA was twice the annual average and the Central Congolian lowland forests where BA was 77% above average, and the Northwestern Congolian lowland forests where BA was 55% above average.

Likewise, several northern regions of Angola experienced record BA (**Figure 2, Figure 3, Table 2, Table 3**). In northern Africa, Mali, Niger, Chad and Sudan all saw high BA in various states or regions that encompass the semi-arid Sahel region, though these anomalies notably occur against a low baseline in most cases due to the typically sparse vegetation fuel loadings in such regions. **Appendix A (Section A1)** provides a more complete summary of the fire season in Africa based on the regional panel assessment.

2.2.1.5. A Low Fire Year in Eurasia

Asian and European biomes generally experienced a low fire year that contributed towards the below-average global BA total in 2024-25 (**Figure 1, Figure 2**). BA was around 50 thousand km² (71%) below average in the Asian temperate grassland, savannah and shrubland biome, aligning with unusually low number of fires (44% below average) and unusually low fire sizes (95th percentile fire size was 42% below average). BA was also 42 thousands km² (62%) below average in the Asian xeric shrublands, and with fire counts 22% below average and 95th percentile fire sizes 39% below average. Elsewhere, BA was 9 thousand km² (11%) below average in the Asian (sub)tropical moist broadleaf forests, with fire counts and sizes both 10-15% below average. The below-average fire extent in all of these regions translated into below-average C emissions, though not in direct proportion because the combustion of vegetation per unit BA also varied compared with previous years (**Figure 1**). For example, while BA was 11% below average in the Asian (sub)tropical moist broadleaf forests, C emissions were 54% (85 Tg C) below average signifying that areas that did burn tended to do so with anomalously low severity. Across Asia as a whole, the total BA was 99 thousand km² (26%) below average during 2024-25, the 4th lowest annual total on record, and C emissions were 119 Tg C (28%) below average, the 5th lowest on record.

While most regions of Asia experienced a low fire year in general, there were some notable exceptions. Many states of northeast India and Nepal experienced high-ranking or record-breaking levels of BA or C emissions (**Figure 2, Figure 3**), highlighting a coherent regional-scale anomaly during 2024-25. Similarly, in northeast Asia where 2 provinces of China (Heilongjiang and Jilin; **Table 2**), 2 provinces of South Korea, and 7 prefectures of Japan experienced record-breaking BA or C emissions and many neighbouring regions likewise experienced high-ranking fire years (**Figure 2, Figure 3**). **Appendix A (Section A2)** provides a more complete assessment of the fire season in Asia.

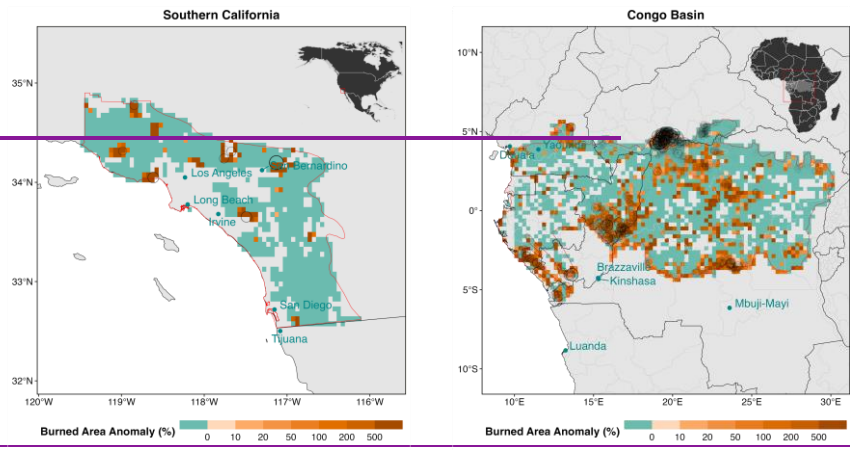
Though less impactful on the global BA and C emissions totals than in the vast Asian biomes, the 2024-25 fire season was notably another low fire year in Europe. For example, BA was 13 thousand km² (59%) below average in the European temperate broadleaf and mixed forests, chiefly because the number of fires was 72% below average. BA was also 12 thousand km² (40%) below average in the European temperate grassland, savannah and shrubland biome, chiefly because the number of fires was 44% below average. Across Europe as a whole, the total BA was 30 thousand km² (49%) below average during 2024-25, the 4th lowest annual total on record, and C emissions were 5 Tg C (22%) below average, the 7th lowest on record.

Despite the low fire activity in Europe, there were several exceptions in southeast Europe. In regions of Serbia, North Macedonia, and western ~~Turkey~~Türkiye experienced record high BA or C emissions in 2024-25. Further north, several eastern regions of Ukraine experienced record-breaking fire C emissions, with some suggesting a link between elevated ignitions and the ongoing conflict in the country (European Commission Joint Research Centre, 2025). **Appendix A (Section A3)** provides a more complete assessment of the fire season in Europe based on regional panel assessment.

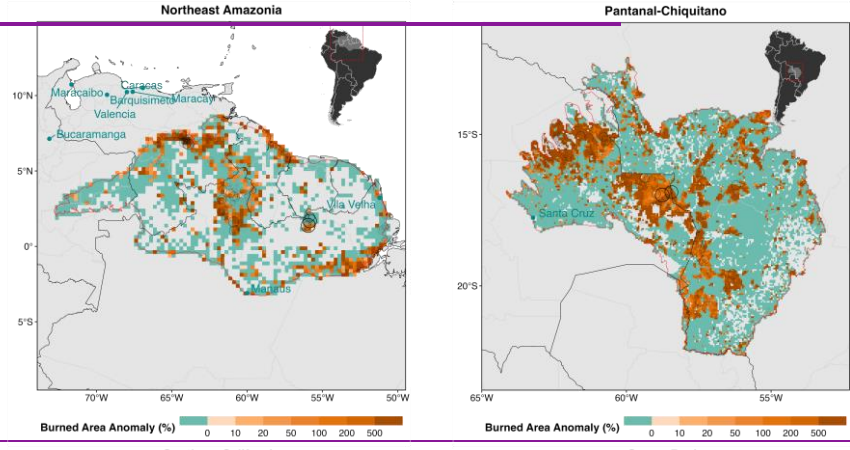
2.2.2. Focal Events of this Report

In this year's report, we identify four focal events with global relevance for further study across **Sections 4-6**. The four events are Northeast Amazonia, the Pantanal-Chiquitano, Southern California, and the Congo Basin (**Figure 5**), and our reasons for selecting these particular events are detailed below. In **Sections 4-6**, our analyses explain the causes of each of the events (**Section 4**), evaluate the predictability of the events (**Section 4**), attribute the events to climate change and land use factors (**Section 5**), and predict the likelihood of similar events under future climate change scenarios (**Section 6**).

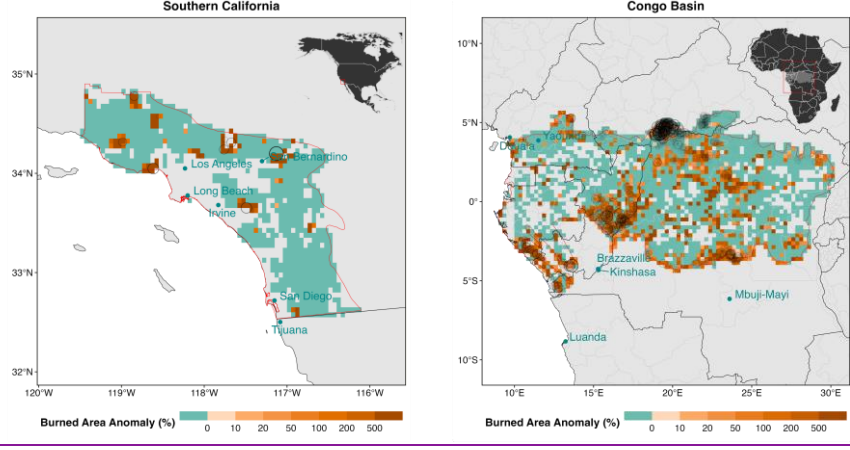
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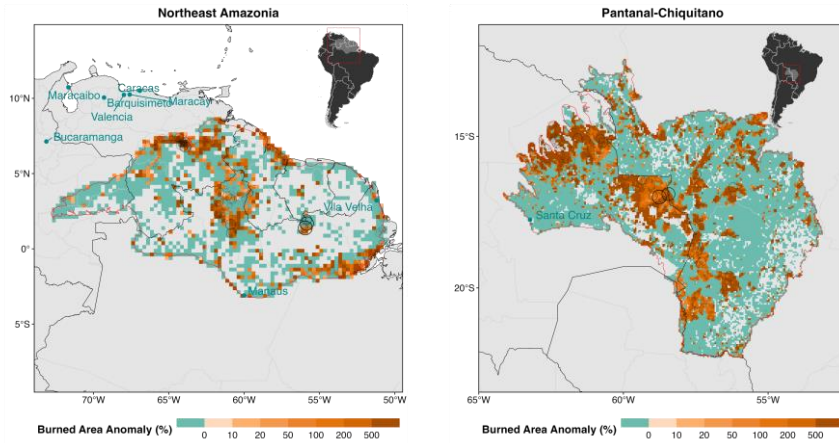


Figure 5: Spatial distribution of burned area (BA) anomalies during 2024-25 relative to the mean annual BA (%). BA is shown at 0.25° resolution (Northeast Amazonia and Congo Basin) or 0.05° resolution (Pantanal and southern California). Fire ignition points (open circles) from the Global Fire Atlas are also shown for the fires with sizes in the upper quartile regionally during 2002-2025, with the largest fires for each region displaying as the largest and most visible circles.

2.2.2.1. Northeast Amazonia (January-March 2024)

2.2.2.1.1. Study Area and Fire Regime

The Northeast Amazonia region here refers to the moist tropical forest ecoregions northeast of the Amazon river and the Rio Negro tributary, mostly including Amazonia but also including the Guianan Shield forests that extend into Venezuela, Guyana, Suriname, and French Guiana (Figure 5). [Fire regimes in northeastern Amazonia reflect interactions among ecosystems, human activity, and climate. Ecological heterogeneity, spanning humid and floodplain forests, natural grasslands, and savanna formations, produces marked variation in fire impacts. Savannas and forest-savanna transition zones in Roraima, Venezuela, and the Guianas are relatively fire-adapted, historically experiencing low-intensity surface burns at multi-year intervals \(Alvarado et al., 2020; Pivello et al., 2021\). Yet their resilience declines under more frequent or intense burning. Fire-sensitive ecosystems such as humid forests and wetlands are even more vulnerable, with elevated fire pressure threatening long-term stability and biodiversity \(Alvarado et al., 2020\). In these ecosystems, anthropogenic climate forcing is most likely driving increases in fire activity. Most ignitions now arise from human use, particularly pasture management, with escaped fires the dominant source of wildfires \(Cano-Crespo et al., 2015\). Burning follows distinct seasonal patterns: August–October in the southern Amazon \(Jakimow et al., 2018\), December–January in coastal savannas \(Brunel et al., 2021\), and January–March in the northeast \(Carvalho et al., 2021\). These dynamics help explain the late peak of the 2024 event. Climatic variability remains the principal temporal control \(Brando et al., 2020; Berenguer et al., 2021\), while land-use practices and socio-economic drivers shape fire use \(Cammelli et al., 2020\). Addressing these coupled drivers requires management that integrates ecological context, community practices, and regional policy frameworks \(Bilbao et al., 2019\).](#)

2.2.2.1.2. Extreme Event

We specifically target the period January-March 2024. In this region, as in other parts of the northern hemisphere tropics, our global March-February fire season definition can be misaligned with local fire seasonality, specifically where fire seasons span two calendar years. Although this event straddles the boundary between the 2023-24 and 2024-25 fire seasons, we include it here to ensure that significant fire activity is not excluded solely due to the constraints of our reporting framework. **Section 2.2.1** discusses the regional anomalies that led this region to be identified (e.g. **Figure 2, Figure 3**), with further review of the fire season provided by our expert panel in **Appendix A (Section A6)**. It emerges as a major event of global relevance for the following reasons:

- **Record-breaking burned area in forests:** The area of forest burned was more than four times (+332%) the average, and the highest on record, while total BA (including non-forests) was also 67% above average. In forests, 8 continuous months of the fire season (March-September 2024) had BA above the climatological mean, peaking in March 2024. The most pronounced anomalies occurred in the Northern Amazonian savannas around Roraima, the forest-savanna transition zones of northern Venezuela and southern Guyana, and the coastal ecosystems near the Guyana-Suriname border (**Figure S5**).
- **Carryover from the previous fire season:** A new record for total BA had been set in the previous fire season (2023-24) mostly due to an anomalously high count of large fires in non-forests. The transition of anomalously high fire activity into forest environments during the 2024-25 fire season was a distinguishing factor.
- **Anomalous fire counts:** The large BA anomalies were explained by an anomalously high number of fires, with 1,500 (52%) more fires than the average fire season.
- **Widespread forest loss:** Highest rates of forest loss (stand-replacing fire extent) since 2016 recorded in Amazonia with 60% attributed to wildfires.
- **Disproportionate impact on rural, traditional populations and Indigenous territories:** Fires degraded air quality and destroyed crops, homes, and native vegetation, intensifying food and water insecurity for those living in the region, including Indigenous peoples. The compounded effects of fire and drought deepened the humanitarian crisis in the Yanomami Territory and local organizations estimate at least 70,000 people across urban and rural areas without access to clean water.

2.2.2.2. *Pantanal and Chiquitano (August-September 2024)*

2.2.2.2.1. Study Area and Fire Regime

The Pantanal-Chiquitano region here refers to the areas draining into the Pantanal (IBGE, 2021), the world's largest tropical wetland area, and the Chiquitano dry forest ecoregion in Bolivia (**Figure 5**).

In the Pantanal, the annual cycle of flooding (October–March) and drying (April–September) plays a central role in shaping fire regimes. During the wet season, extensive inundation keeps most of the landscape too moist to burn. As waters recede, grasses and savanna vegetation dry out, creating windows of flammability, but under normal conditions fires remain patchy and largely restricted to grasslands, savannas and wetland margins (Damasceno-Junior et al., 2021). When this cycle is disrupted, often by multi-year droughts, large areas stay dry for longer, exposing grasslands, forests, and even peat-rich soils to extensive burning. In the Chiquitano, surface fires are the dominant type, frequently originating in deforested or agricultural areas before spreading into forest edges. Human activity is the primary ignition source in both regions (Romero-Muñoz et al., 2019; Menezes et al., 2022), with escaped agricultural and pasture-renewal fires driving many of the catastrophic events of recent years. These coupled processes illustrate that fire activity in the Pantanal and Chiquitano is no longer governed by climate alone but increasingly by altered hydrology, land-use frontiers, and the intensity of human fire use (Barbosa et al. 2022).

2.2.2.2.2. Extreme Event

We specifically target the period January-March 2024, when the most substantial anomalies in BA were observed (Figure S6). Section 2.2.1 discusses the regional anomalies in Brazil and Bolivia that led this region to be identified (e.g. Figure 2, Figure 3), with further review of the fire season provided by our expert panel in Appendix A (Section A6). It emerges as a major event of global relevance for the following reasons:

- **Record-Breaking burned area:** BA in the Pantanal-Chiquitano region was almost triple (+196%) the annual average, and the highest on record. This anomaly included a +466% BA anomaly in forests. There were 8 continuous months (March-October) with BA above the climatological mean, oriented around a peak in August 2024.
- **Record-Breaking carbon emissions:** Fire C emissions were 6 times (+502%) the annual mean, driven up by the large anomaly in forest fire C emissions in the period.
- **Record fire size and spread:** The 95th percentile fire size for the region was over three times (+226%) the average and the 95th percentile rate of growth was 88% above average, signifying that large, fast-spreading fires drove up the anomalous BA total in the region.
- **Severe air quality degradation:** Over 900 µg/m³ of fine particulate matter (PM_{2.5}) was recorded in September 2024, which is 60 times above WHO standard.
- **Economic losses:** Agribusiness losses due to wildfires reached R\$ 1.2 billion (~US\$222 million) in the Pantanal, the biome's main economic sector.
- **Challenges in response:** 78 days of firefighting effort which involved multiple actors was marked by significant access and logistical challenges in remote regions, making it difficult to reach and support isolated communities.

2.2.2.3. Southern California (January 2025)

2.2.2.3.1. Study Area and Fire Regime

Southern California here refers to the Mediterranean portions of seven counties in California (Los Angeles, Orange, Riverside, San Bernardino, San Diego, Santa Barbara, and Ventura; Figure 5). The Mediterranean portions are defined based on the ecoregional definition of the US Environmental Protection Agency (EPA, 2024).

Southern California's fire regime reflects the interaction of Mediterranean-climate vegetation, frequent extreme weather, and dense human presence. The region is dominated at lower elevations by chaparral shrublands that historically experienced fire return intervals of several decades, and frequent-fire forests at higher elevations. Southern California's chaparral shrubland ecosystem is distinctly different from the frequent-fire dry forests elsewhere in the western US, where over-suppression is often discussed as a driver of extreme burning (Keeley and Fotheringham 2001). Over-suppression is less likely to have influenced the January 2025 fires. Pre-colonial Indigenous burning shaped fire patterns and fuel distribution (Keeley, 2002), but Euro-American settlement and ongoing urban expansion have altered ignition patterns and increased the frequency of fire (Keeley et al. 1999). Fire suppression resources are extensive but overwhelmed during extreme fire weather, particularly where land use and climate 'whiplash' between extremely wet and dry years produce heavy fuel loads (Swain et al. 2025). Consequently, the contemporary fire regime is characterized by highly variable burned area, frequent wildland-urban interface impacts, and substantial sensitivity to meteorological extremes such as katabatic 'Santa Ana' winds. This context highlights that while human activity strongly influences fire occurrence and exposure, the underlying bioclimatic and ecological conditions continue to govern fire behaviour in southern California (Jin et al. 2015, Parks et al., 2015).

2.2.2.3.2. Extreme Event

Although California as a whole did not experience a particularly strong fire season in 2024-25 from the vantage of BA or fire C emissions (e.g. **Figure 2, Figure 3**), the regional expert panel identified the numerous wildfires affecting LA and surrounding counties in January 2025 as a major event of the 2024-25 fire season (see **Appendix A Section A4**), with the Palisades and Eaton fires in particular leading to loss and damage in the suburbs of LA. We specifically target the period January 2025 when the most substantial anomalies were observed (**Figure S7**). Southern California emerges as a major event of global relevance for the following reasons:

- **High fatalities and structure loss.** Over 11,500~~750~~ homes were destroyed across Los Angeles County, and at least ~~3031~~ lives were lost (Los Angeles County Coroner, 2025; Wikipedia, 2025). The Palisades Fire severely damaged or destroyed nearly 8,000 structures~~5,614 homes~~, while the Eaton Fire impacted over 10,000 severely damaged or destroyed 6,150 homes and hundreds more commercial structures (CALFIRE, 2025; Wikipedia, 2025). Notably, most of the destroyed structures were in moderate-to-high density suburban and urban neighborhoods, not on large lots in rural, low-density areas as has historically been the case for structure loss wildfires in the US.
- **Mass evacuations.** At least 153,000 people evacuated, with up to 200,000 under evacuation warnings or orders during the peak of the crisis (USGS, 2025b; NPR, 2025; Wikipedia, 2025).
- **Air quality impacts.** Air and municipal water quality were heavily impacted by the fires, contributing to negative health outcomes for thousands. During the fires, peak PM_{2.5} levels were recorded at 483µg/m³ (an order of magnitude greater than the 35 µg/m³ daily standard set by the US Environmental Protection Association), part of a prolonged period of Hazardous air quality (California Air Resources Board, 2025). Over 400 excess deaths in Los Angeles County have since been attributed to exposure to poor air quality during January 2025 (Paglino et al., 2025).
- **Water quality impacts.** Municipal water supplies were considered unsafe for several weeks following the fires for tens of thousands of residents in the affected areas (City of Pasadena, 2025). ~~As a political~~ response to the fires ~~outside Los Angeles~~, over 8.3 million cubic meters of water were released from federal reservoirs over 200km away from Los Angeles in central California, a move which has been criticised because this water did not supply southern California, happened well after the fires were controlled, and because it would otherwise have been used for irrigation in the Central Valley (Levin et al., 2025).
- **Exceptional economic loss.** Total economic losses were estimated at US\$140B including property destruction, health costs, business disruption, and infrastructure impacts, making this one of the most costly wildfire events in US history (LAEDC, 2025; UCLA Anderson School of Management, 2025).
- **Wider economic disruption.** The fires are projected to cause US\$4.6-8.9 billion in lost economic output over five years, 25,000-50,000 job-years lost, and labour income reductions of US\$1.9-3.7 billion (LAEDC, 2025). The Palisades and Eaton fires affected almost 2,000 businesses (LAEDC, 2025). As LA is also the largest port on the US Pacific coast, the fires impacted broader supply chains that run through the port of LA (Terrill, 2025).
- **High insured losses.** Industry estimates have placed insured losses in the range of to US\$20-75 billion (Li and Yu, 2025; Morningstar DBRS, 2025; Insurance Insider, 2025), placing substantial additional stress on the already volatile home insurance market in California and on most global reinsurers.
- **Housing and affordability crisis.** Thousands of affordable housing units were destroyed, worsening Southern California's housing shortage, displacing large numbers of lower-income residents, and exacerbating the problem of homelessness in the region (Mattson-Teig, 2025; Li and Yu, 2025; Booth, 2025). This triggered a ripple mass displacement into both surrounding communities and beyond in the months following the fires (NYT, 2025).

- **Debris flows.** The geology of southern California is highly conducive to erosion and debris flows after wildfires. Several debris flows following high-intensity rainfall events in the weeks after the fire produced further damage and required hundreds of additional evacuations in and near the affected areas (USGS, 2025a).

The fires in Southern California have already been subject to several detailed investigations, which found that the fires were driven by exceptionally late onset of winter rains that extended the fire season into January, unseasonably warm winter temperatures, fuel buildup from very wet conditions in the prior ~~year to two years~~, and powerful Santa Ana winds locally exceeding 130 km/h, creating extreme fire weather conditions that propelled fires to progress downhill from wildlands into the built environment and become an urban conflagration (Barnes et al., 2025; Garrett, 2025). The potential for extreme wildfires to develop under dry downslope winds was predicted several days in advance, including by the National Interagency Fire Center (NIFC), the National Weather Service (NWS), and the Storm Prediction Center (SPC; see summary by Wikipedia, 2025) as well as by specialist commentators (e.g. Swain, 2025).

2.2.2.4. **Congo Basin (July-August 2024)**

2.2.2.4.1. Study Area and Fire Regime

The Congo Basin region here refers to the moist tropical forest ecoregions of equatorial Africa (**Figure 5**). Here, fires have historically been rare because short dry seasons and high moisture constrain fuel availability and limit the natural ignitions and spread of fires (Wimberly et al., 2024). Interannual fire variability is positively correlated with higher temperatures and atmospheric drying, and widespread outbreaks were recorded under the anomalously warm and dry conditions of some El Niño events, such as in 2015-2016 (Wimberly et al., 2024; Dwomoh et al., 2019; Verhegghen et al., 2016). Recent satellite observations demonstrate increasing fire occurrence across multiple Congolian ecoregions, particularly in the central lowland and swamp forests, where active fire detections approximately doubled between 2003 and 2021 (Wimberly et al., 2024). These trends are closely associated with deforestation and fragmentation in the central and western parts of the basin, which are largely driven by small-scale agriculture and logging (Shapiro et al., 2021, 2023). Land-use change alters canopy structure and understory microclimates, increasing the likelihood that anthropogenic ignitions will spread into forested areas (Zhao et al., 2021; Dwomoh & Wimberly, 2017). The contemporary fire regime is thus characterised by rising exposure of tropical forests to anthropogenic ignitions, heightened sensitivity to climate extremes, and growing implications for carbon storage, biodiversity, and local livelihoods (Wimberly et al., 2024).

2.2.2.4.2. Extreme Event

We specifically target the period July-August 2024, when the most substantial anomalies in BA were observed (Figure S8). Section 2.2.1 discusses the regional anomalies that led this region to be identified (e.g. Figure 2, Figure 3), with further review of the fire season provided by our expert panel in Appendix A (Section A6). #The Congo basin emerges as a major event of global relevance for the following reasons:

- **Record-breaking burned area:** Highest-ranked BA on record at 28% above the annual mean due to there being 4,000 (20%) more fires than in the average year. There were 7 continuous months with BA above the climatological mean. For seven months in a row, the observed burned area was greater than the historical average for those same months, based on the reference climatology since 2002. The largest fire anomalies were observed during July and August (**Figure S8**), especially in southern Democratic Republic of the Congo, northern Angola, and parts of the Republic of the Congo.
- **Unprecedented role of fire in primary forest loss:** Forest loss statistics from the recent Global Forest Watch report (Goldman et al., 2025) showed that wildfires were

the dominant driver of a more than doubling (+150%) of rates of forest loss in the Republic of the Congo and the Democratic Republic of the Congo during 2024 versus 2023, representing the highest rates of primary forest loss since 2015.

- **Sparse reporting and poor media coverage:** Reporting on the occurrence, drivers, and consequences of fire is extremely sparse in this region, including by government agencies and the international and national news media. This demonstrates that extreme fire events in this region are often overlooked, making it an intriguing important case study to investigate in this report.

3. Impact Assessments

In this edition of the report, we introduce new routine regional assessments of fire impacts on society in terms of population exposure to fire, physical asset exposure to fire, the exposure of carbon projects to fire, and the degradation of air quality through emissions of fine particulate matter (PM_{2.5}). For our air quality analysis, estimates are generated for the focal events only (**Section 2.2.2**). In all other cases, estimates are provided for each of the regional layers detailed in **Table 1**, mirroring our approach to providing regional summaries of BA, C emissions, and individual fire properties (**Section 2.1.2**).

3.1. Methods

3.1.1. Population Exposure Assessment

Population exposure estimates are produced using the global risk assessment platform CLIMADA (Aznar-Siguan and Bresch 2019). CLIMADA has previously been validated and applied to systematically quantify exposed population to a variety of natural hazards globally, such as river floods (Kam et al. 2021) and tropical cyclones (Stalhandske et al., 2024; Kam et al. 2024). The BA hazard set is set up using the MCD64A1 MODIS BA product (Giglio et al. 2018). The original BA data are aggregated monthly on a regular grid with a resolution of 150 arcsec and expressed as the fraction of total cell area burned. For the spatial distribution of exposed population, we use Gridded Population of the World (Doxsey-Whitfield et al., 2015), which is spatially reaggregated on the same grid as the hazard using the LitPop exposure layer (Eberenz et al. 2020). The population exposed to wildfires is estimated by multiplying the BA fraction (BA expressed as a fraction of burnable area) of each cell by the population present in each grid cell. As a complementary approximation to the main analysis, a single displacement share is derived by comparing population exposure estimates with reported displacement figures from the Internal Displacement Monitoring Center (IDMC, 2025), acknowledging that exposure only partially translates into impact. Event records are matched to BA observations following the methodology described in Riedel et al. (2025). We compute the ratio between recorded impacts and exposed values for each event and provide the median of these damage ratios across events.

The data produced using these methods are available from Steinmann et al. (2025).

3.1.2. Physical Asset Exposure Assessment

Physical asset exposure estimates are produced using the global risk assessment platform CLIMADA (Aznar-Siguan and Bresch 2019). CLIMADA has previously been validated and applied to systematically quantify economic impacts resulting from exposure of physical assets to a variety of natural hazards globally (Stalhandske et al. 2024), including fires (Lüthi et al. 2021). The exposure layer LitPop (Eberenz et al. 2020) was used to spatially distribute national-scale macroeconomic indicators as a function of night light intensity (Román et al. 2018) and population density (Doxsey-Whitfield et al., 2015) within national geographical domains. We disaggregate country-based produced capital estimates (World Bank, 2024c) for

the year 2018 to approximate physical asset density in US dollars (US\$). Physical asset exposure to wildfires is estimated by multiplying the BA fraction of each cell by the physical asset totals present in each grid cell (analogous to our analysis of population exposure, **Section 3.1.1**). In addition to this analysis, a single overall loss fraction is provided recognising that exposure tends to overstate actual asset damage. This fraction is derived by comparing modelled exposure estimates with asset damages from wildfire events, as reported in the Emergency Events Database (EM-DAT; Delforge et al. 2025) maintained by the Centre for Research on the Epidemiology of Disasters (CRED). Event records are matched to BA observations following the methodology described in Riedel et al. (2025). We compute the ratio between recorded impacts and exposed values for each event and provide the median of these damage ratios across events.

The data produced using these methods are available from Steinmann et al. (2025).

3.1.3. Carbon Projects Exposure

We estimated the exposure of carbon offset projects to fire by combining a large set ($n=927$) of project boundaries for forestry projects in Latin America ($n=394$), northern America ($n=316$), Eurasia ($n=150$), Africa ($n=60$), and Australasia (7) with information on fire and climate. Project boundaries were sourced from BeZero Carbon Ltd., who have collated and digitised boundaries for all nature-based projects in the Voluntary Carbon Market (VCM). Information on annual BA was derived from the MCD64A1 collection 6.1 data (Giglio et al., 2018) and this was combined with information on land cover from MCD12Q1 collection 6.1 (Sulla-Menashe et al., 2019) to separate forest from non-forest fires. To evaluate drought conditions, we calculated the 12-month Standardized Precipitation Evapotranspiration Index (SPEI) using data from ERA5-Land (Muñoz-Sabater et al., 2021) calibrated over the 1980-2014 period.

We evaluated fire activity during the 2024 calendar year in the context of long-term trends in drought and fire risk. First, to assess how 2024 compared to previous years since 2001, we calculated the number of carbon projects affected by fire in each year and the average percentage of project area burned per year (%). Second, to place this in the context of climate change, we calculated the 2024 drought anomaly as the 2024 SPEI minus the long-term average SPEI (1980-2023).

3.1.4. Air Quality Impact Assessment

The human health risks associated with fire smoke pollution are well established. Smoke contains a toxic mix of gases, compounds, including ozone and carbon monoxide, and volatile organic compounds (VOCs), as well as fine particulate matter ($PM_{2.5}$) that can carry heavy metals and environmentally persistent free radicals (Hamilton et al., 2021; Andreae, 2019; Fang et al., 2023). Even short-term exposure to these pollutants has been associated with increased risk of cardiovascular and respiratory illnesses, including asthma exacerbation, reduced lung function, and acute infections (Johnston et al., 2021; Xu et al., 2024; Chen et al., 2021; Xu et al., 2023; Aguilera et al., 2021; Zhang et al., 2025). Furthermore, wildfire smoke contributes to increased mortality, particularly among vulnerable populations. In addition to these physiological effects, heavy smoke can significantly reduce visibility, compounding health risks by increasing the likelihood of injuries during regular driving, evacuation, or emergency response (Gill and Britz-McKibbin, 2020), and generates lasting mental health effects amongst exposed or displaced communities (Humphreys et al., 2022).

To quantify the contribution of fires to degraded air quality we used the global model framework utilised by the Copernicus Atmosphere Monitoring Service (CAMS) to simulate concentrations of fine ($<2.5 \mu m$ diameter) particulate matter ($PM_{2.5}$; Peuch et al., 2022). One

of the key objectives of CAMS is to monitor and forecast global atmospheric composition including smoke from vegetation fires. ~~Fires~~Fire emissions in CAMS ~~are prescribed by~~derive from the Global Fire Assimilation System (GFAS; Kaiser et al., 2012), which calculates hourly estimates of biomass burning emissions by assimilating fire radiative power (FRP) observations from satellite-based sensors and by means of land cover-dependent conversion (FRP to dry matter) and emission factors (dry matter to emitted gas or aerosol species per biome) describing the rate at which about 40 smoke constituents are released into the atmosphere. This study uses GFAS version 1.4, which is the version used currently for the NRT production input of CAMS global and regional forecast services, plus some improvements that include the use of VIIRS FRP retrievals. Spurious FRP observations of no vegetation fire origin are filtered out in GFAS with a static map. GFAS ingests active fire information together with a characterization of its uncertainty, including an uncertainty component related to the satellite sensor detection limit and a solution for partial observational cloud coverage.

Simulations are run with the Integrated Forecasting System extended with modules of atmospheric composition (IFS-COMPO), which describe source, sink, and transport processes of the main reactive trace gases (Flemming et al., 2015; Huijnen et al., 2016) and aerosol species (Morcrette et al., 2009; Remy et al., 2022, 2024) and which, together with satellite observations, is at the core of the CAMS system for the global domain. Mass fluxes of atmospheric constituents from the surface into the atmosphere are either prescribed from CAMS pre-compiled emissions inventories, with some aspects of on-line simulated temporal variability, as in the case of pollutants from the burning of fossil fuels for transportation and electricity, or estimated online at every time step in the IFS when strongly dependent on meteorological conditions as in the case of desert dust and sea salt aerosol and of biogenic fluxes of CO₂. The resolution used is the current operational resolution of 40 km, with 137 vertical levels. GFAS biomass burning emissions are estimated at 0.1° resolution based on FRP observations from the MODIS sensor on both the Terra and Aqua satellites (Giglio et al., 2016) and from the VIIRS sensor on the Suomi NPP satellite (Csiszar et al., 2014). The vertical distribution of fire emissions within the simulation follows the GFAS IS4FIRES injection height estimation (Sofiev et al., 2012; Remy et al., 2017).

To isolate the contribution of extreme fire events to atmospheric PM_{2.5} concentrations, two sets of forecast experiments are run for specific focal events using a similar assessment framework. In the first (“with local fires”), all emission sources of PM_{2.5} were considered including those of anthropogenic, dust, biogenic and other natural origin. In the second (“no local fires”), biomass burning emissions from within the focal event are excluded. The difference in simulated PM_{2.5} concentrations between the two runs then represents the fire contribution to PM_{2.5} within the region.

After PM_{2.5} concentrations had been simulated at a 3-hourly temporal, and 40 km spatial, resolution, we summarised the influence of fires in the region to a daily population-weighted mean PM_{2.5} concentration at ground-level (in units of µg/m³) for each focal region. Population data for the year 2020 from the Gridded Population of the World (GPW) dataset version 4 (Doxsey-Whitfield et al., 2015) were used to weight the values of PM_{2.5} concentration in each grid cell of the focal regions, producing a weighted mean value for PM_{2.5} concentration for each simulated date. This process was repeated for each simulation (“with local fires” and “no local fires”), and daily differences between the simulations were used to assess the additional number of days with poor air quality caused by fires in the focal regions.

To illustrate the scale and intensity of wildfire smoke health-relevant exposure within the 2024-2025 fire season, total population-weighted PM_{2.5} and the isolated contribution of fires to population-weighted PM_{2.5} in a focal region are compared against the World Health Organisation 24-hour mean (15 µg/m³) standard for daily PM_{2.5} exposure (WHO, 2021).

3.2. Results

3.2.1. Population Exposure

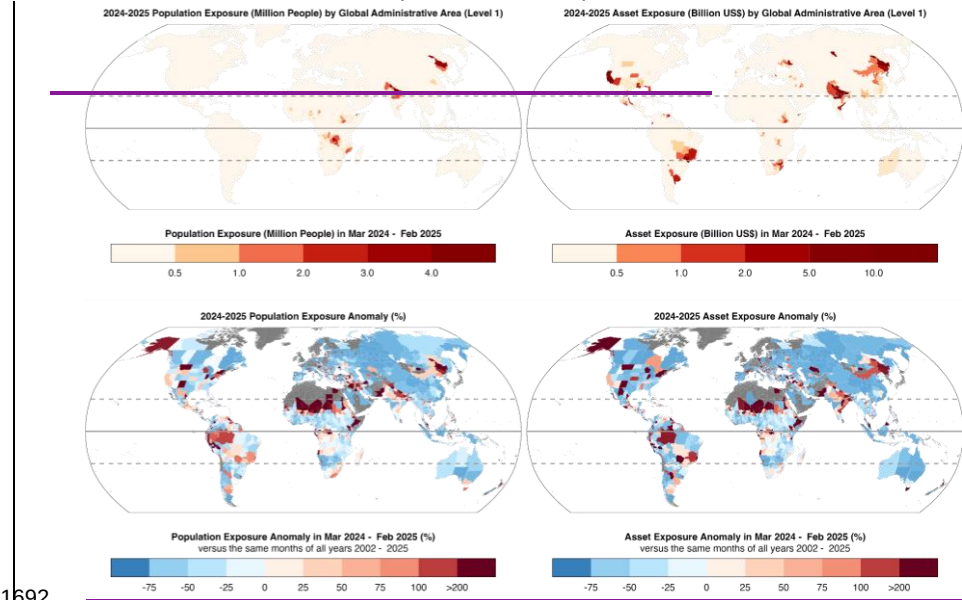
During the 2024-25 fire season, we estimate approximately 100 million people to have been exposed to wildfires worldwide. Exposure was most pronounced across South and Southeast Asia, as well as Central and East Africa. At the country level, India and the Democratic Republic of the Congo show the highest numbers, with around 15 million people affected in each (**Figure 6; Figure S9**). Nigeria, China, Mozambique, and South Sudan also were also exposed substantially, each with more than 5 million people affected. At the subnational level, we estimate the highest population exposures in Uttar Pradesh State (India) with over 4.6 million people, Heilongjiang Province (China) with 3.7 million, and Punjab State (India) with 3.6 million exposed (**Figure 6; Figure S9**). Several provinces in the Democratic Republic of the Congo also exceed 2 million, illustrating how national-level exposure is often driven by a few highly affected administrative regions.

Some of the countries with the most extreme anomalies in fire BA and C emissions, most notably Bolivia, Brazil, and Canada, accounted for only a small share of absolute global population exposure, and showed negative (Canada) to modest positive (e.g. Bolivia and Brazil) anomalies (**Figure 6; Figure S9**). This decoupling highlights the relevance of the spatial distributions of both BA and population to population exposure, which might be low when extensive fires occur in remote places.

Several of the countries with the highest absolute exposures, such as in India and the Democratic Republic of the Congo, showed negative anomalies on a national level, related to the fact that fire-related population exposure in these regions is more recurrent. Nonetheless, on a subnational level, some regions of these countries show considerable positive anomalies, such as in India's Uttar Pradesh State where 4.6 million people were exposed (146% above average) and about half a million being exposed in the provinces Kasai-Central (+33%) and Kongo-Central (+27%) in the DRC. We note that our analysis of population exposure to fire captures exposures to all fire types (see Section 3.1.1), including wildland fires but also fires observed in agricultural settings (e.g. crop stubble removal or for pasture maintenance) or in shifting cultivation systems, which generally pose low direct risks of fatality or injury. The use of fire for agricultural burning is widely documented in Uttar Pradesh (Shyamsundar et al., 2019) and shifting cultivation is a widespread practice in the Congo Basin (Molinario et al., 2015; Tyukavina et al., 2018), with these modes of fire use dominating over other uses in the respective regions (Millington et al., 2022). As a result, some regions ranked highly for population exposure in our analysis are particularly susceptible to inflated estimates due to the prevalence of exposure to relatively low-risk fire types. Recent work has begun to address this issue by quantifying exposure specifically to higher-intensity fires (Teymoor Seydi et al., 2025), an approach we intend to adopt in future editions of this report.

Population exposure anomalies were also high in relative terms across parts of the Middle East and the Balkans (e.g. Jordan, Iran, Iraq, North Macedonia, Albania), the Andes region (e.g., Peru, Ecuador), the Northern coast of South America (Venezuela, Guyana, Suriname; broadly encompassing our focal region of Northeast Amazonia), and Central Sahel (e.g., Niger), as well as isolated cases such as Nepal and Iceland. For example, Jordan shows divergent anomalies in population exposure (+201%) resulting from large subnational regions of Balqa (+322%) and Irbid (+393%). These patterns of exposure mostly align with patterns in BA and carbon emissions (**Section 2**), in the Middle East and the the Balkans, Andes, Northern coast of South America and Central Sahel. Although the absolute number of people affected in some of these countries remains low, the relative anomaly marks a sharp departure from historical patterns.

1676
1677 It is important to distinguish between the exposed and affected population. Based on 521
1678 events in the years 2008-2025 recorded by IDMC (2025), we estimate the damage ratio of exposed
1679 to displaced population to amount to 3.0%. While nearly 100 million people were exposed to
1680 wildfire activity in the 2024-25 season, only a small fraction - 20,046 people (IDMC, 2025) -
1681 were formally displaced (0.02%). Note, however, that this figure likely understates the true
1682 scale of disruption, as displacement records are incomplete. Many affected individuals may
1683 not be forced to leave their homes but still experience substantial short- and long-term
1684 consequences, including health burdens (Gould et al., 2024) and financial distress such as
1685 short-term earning disruptions (Borgschulte et al., 2024), increased missed mortgage
1686 payments (Ho et al., 2023), declines in property values (Huang and Skidmore, 2024), and
1687 lasting reductions in income later in life (Meier et al., 2025). Moreover, recent cases have
1688 emphasised that the number of people impacted by wildfire smoke can be many times higher
1689 than the number of people directly exposed to fire (Jones et al., 2024b; Kolden et al., 2024,
1690 2025; Johnston et al., 2021). As such, these records should be viewed as a conservative lower
1691 bound on the broader human impacts of wildfire exposure.



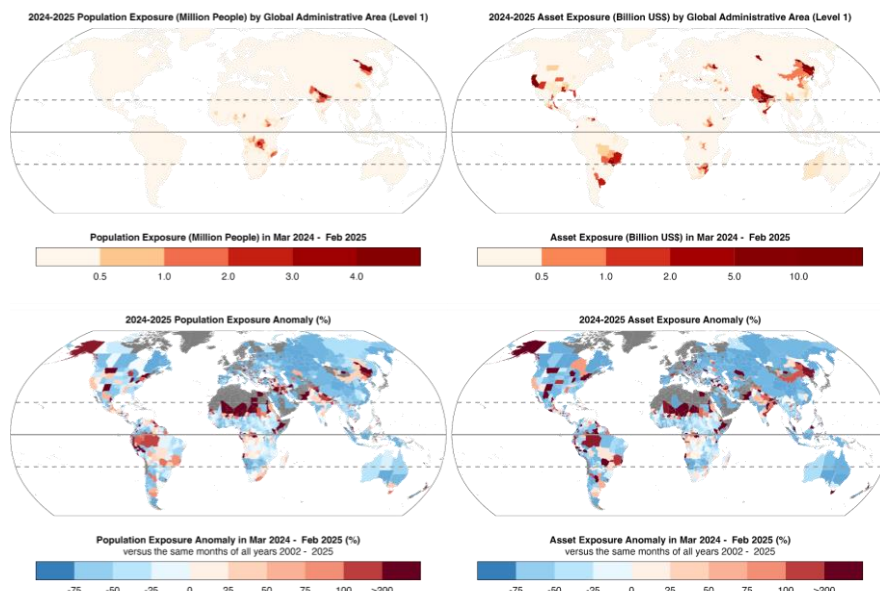


Figure 6: (left panels) Population and (right panels) physical assets exposed to burned area (BA) during the 2024-25 global fire season. The figure shows (top panels) the number of people or the asset value (billion US\$) exposed to fire and (bottom panels) the relative anomaly versus all years since 2002. Results are shown at the national scale in Figure S9.

3.2.2. Physical Asset Exposure

We estimate that physical assets exposed to wildfires during the 2024-25 season amounted to US\$215 billion worldwide. The highest asset exposures were concentrated in a mix of middle- and high-income countries, led by India (US\$44 billion), the United States (US\$26 billion) and China (US\$17 billion), followed by Venezuela, South Africa, and Brazil (Figure 6; Figure S9). While India, and to a lesser extent Brazil and China, ranked highly in both population and asset exposure, the asset exposure landscape broadens to include developed countries such as the United States and South Africa (US\$14 billion). This divergence not only reveals different spatial patterns of wealth and infrastructure but also the concentration of high-value assets in certain subnational regions (Figure 6; Figure S9). For instance, South Africa's Gauteng province, its economic hub, ranked among the most exposed globally at US\$8 billion, despite the country's moderate population exposure. Similarly, in the United States, California alone accounted for over US\$17 billion in exposed assets, driven largely by the severe January 2025 wildfires (US\$14 billion) discussed in Section 2.2.3. These estimates are still low in comparison to damage records provided by EM-DAT for the LA fires (US\$52.5 billion). This difference is likely caused by an underestimation of the affected exposure, which consisted of exceptionally high-value structures not represented by LitPop. This also explains the underestimation of the asset exposure anomaly in California, which is less pronounced (+60%) than in other states and regions of the world (Figure 6).

In contrast, Central African countries such as the Democratic Republic of the Congo, Nigeria, and Mozambique, which featured prominently in population exposure, did not rank highly in terms of exposed assets (Figure 6; Figure S9). The exception is South Sudan (US\$4 billion), where asset exposure remains substantial. The data also highlights high absolute asset exposure in Mexico, TurkeyTürkiye, and the Russian Federation, each with national totals

around US\$8 billion. At the subnational level, exposure was concentrated in economically important regions, including Izmir in [TurkeyTürkiye](#) (US\$3 billion), Mexico City (Distrito Federal; US\$3 billion), and Russia's Kemerovo and Rostov regions (approximately US\$3 billion each). These patterns underscore how wildfire-related asset exposure is shaped by the intersection of fire occurrence with concentrated infrastructure and economic activity.

Asset exposure anomalies for the 2024-25 fire season, expressed relative to the same months of all previous fire seasons from 2002-2024 ($n = 23$), reveal several hotspots with unusually high physical asset exposure. Notable positive national-level anomalies were concentrated across the Middle East (e.g., Iraq, Syria), Southeast Europe and the Balkans (e.g., Albania, Bosnia and Herzegovina, Greece), parts of the Sahel and Horn of Africa (e.g., Niger, Eritrea), and the northern tropical [parts](#) of South America (e.g., Ecuador, Colombia, Guyana) (**Figure 6; Figure S9**). At the subnational level (**Figure 6; Figure S9**), pronounced relative anomalies were observed in regions not necessarily among the highest in absolute asset exposure. For example, many of the strongest asset exposure anomalies were highly localised, including regions of Chad, Sudan, Brazil, and Pakistan, where this season's values sharply deviate from past levels (**Figure 6; Figure S9**). In contrast, while California recorded the highest total asset exposure, its relative anomaly was modest, reflecting its regular exposure to fire. These spatial contrasts underscore that extreme fire seasons can affect both high-value regions and those with historically lower risk.

A comparison between asset exposure anomalies and BA anomalies (**Figure 6**) shows areas of both alignment and divergence. Overlaps are evident in Venezuela, western Brazil, Niger, and parts of India and Bolivia, where elevated fire activity coincided with high asset exposure. In contrast, strong BA anomalies in parts of equatorial Africa and Russia were not matched by anomalous asset exposure. This disconnect underscores that fire activity alone is not a sufficient proxy for physical asset impact. Rather, extensive burns in remote or forested areas may have limited consequences for built infrastructure, whereas smaller fires near wildland-urban interfaces can generate disproportionately high asset exposure (Calkin et al., 2023).

As with population exposure, asset exposure does not equate to realised impact. Comparing modelled exposed assets with reported EM-DAT figures, economic losses from 105 historic wildfire events in the time period 2002-2025 show a damage ratio of around 29% of exposed asset value. While a modelled US\$215 billion in physical assets were exposed to wildfires in 2024-25, a smaller sum of US\$57 billion in realised damages was recorded by EM-DAT, or around one-quarter of our exposure estimate. Note, that these figures reflect differences in scope and data quality. EM-DAT's total economic damage records may include indirect losses, such as business interruption and sectoral impacts. Its definition is broad, source-dependent, and rarely disaggregated. Thus, reporting is uneven and regionally biased due to variation in local capacity and data availability (Mazhin et al., 2021, Jones et al., 2023). In contrast, our modelled asset exposure offers a consistent estimate of physical assets at risk, representing the maximum potential asset loss. Yet, it does not represent realised or total economic damage. While both measures have limitations, together they help to characterise the scale of global wildfire-related economic impacts.

3.2.3. Carbon Projects Exposure

Forestry projects can provide cost-effective climate mitigation and co-benefits to society and biodiversity, though their outcomes depend on complex interactions between project activities and their local ecological and social context (Holl and Brancalion, 2020). Wildfires present a growing threat to forest carbon offset projects, posing risks to the permanence of stored carbon (Anderegg et al., 2020) and thus credit integrity (Badgley et al., 2022) and the financial viability of project activities (Conte and Kotchen 2010, Michaelowa et al., 2021). Forestry projects can focus on emissions avoidance (e.g. REDD+), emissions removal (e.g. afforestation or forest restoration), or a combination (e.g. improved forest management). Here

we evaluate fire activity during the 2024 calendar year across an unprecedented number of forestry projects in the Voluntary Carbon Market (VCM), and place results in the context of long term trends in fire risk.

The 2024 fire season was characterized by anomalously high fire activity across the 927 projects evaluated. In total 169, or 18% of projects recorded BA in 2024, a record over the observational period (2001-2024) (Figure S10 (a)). This coincided with record annual BA with 1.6% of project areas affected on average (Figure S10 (b)). Regional drought extremes were likely responsible for the observed uptick in fire activity during 2024, with drought conditions in 72% of projects exceeding the long-term (1980-2023) average and, in 13% of projects, exceeding extreme (SPEI < -2) drought conditions (Figure S10 (c)).

Interestingly, observed anomalies vary regionally and further depend on project activities. Exceptional drought conditions in Latin America resulted in a record number of projects being affected by fire but total BA was just short of previous peak years. In this region, many projects focus on the avoidance of deforestation (38%), and in addition to climate, fire risk is driven by changing land cover and land use over time (Alencar et al., 2015). In comparison, in northern America a smaller number of projects are prone to fire annually and the majority (93%) of projects focus on improved forest management. Here, a record average ~~burned area~~BA was observed, but the total number of projects affected was modest and aligned with average drought conditions. Africa had the highest average BA but 2024 was a low fire year, aligned with long term BA trends in African savannas and woodlands (Andela et al., 2014), and a relatively large number of projects focused on afforestation or forest restoration (52%), which may result in decreasing fire activity over time.

Notably, despite increasing fire risk, about 46% of projects did not experience any BA over the observational period, and 67% of projects were at moderately low risk from fire (with less than 0.5% burned annually in the forests within a 50-km buffer zone around the project).

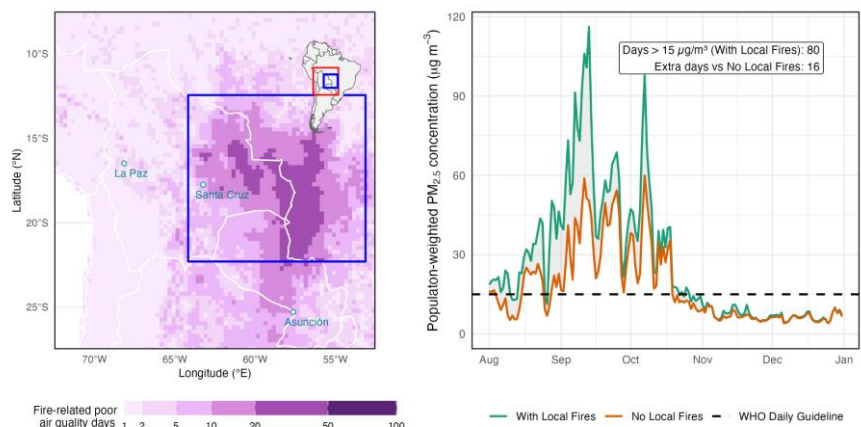
Aligned with long-term changes in fire weather (Jolly et al., 2015; Abatzoglou et al., 2019), we found that the majority of forest carbon projects faced anomalous drought conditions in 2024. The 2024 fire season affected a record number of forest carbon projects globally, resulting in an unprecedented annual percentage of BA within project boundaries. High-integrity forest carbon projects can help to mitigate global climate change, and ~~we find some evidence prior work has shown~~ are also reducing can reduce fire risk locally. (Croker et al., 2023). Nonetheless, the quality of carbon credits issued by nature-based projects depends on the permanence of the carbon emissions avoided or removed, which we show to be increasingly at risk.

3.2.4. Air Quality Impact

Here, we present estimates of the concentration of fine particulate matter (PM_{2.5}) that the average person in the Pantanal-Chiquitano region experienced due to wildfire smoke emissions (the population-weighted PM_{2.5} concentration; **Figure 7**). In the Pantanal-Chiquitano, the population-weighted PM_{2.5} concentration exceeded the WHO daily PM_{2.5} daily standard of 15 µg/m³ on most days from August to November (**Figure 7**), with only 30 days between July and October falling below the threshold, most of which were in early July. Considering fire emissions alone, the average person experienced PM_{2.5} above 15 µg/m³ on 16 additional days between July to October due to local fire emissions, which is slightly lower (20%) than previous 30% estimate of the contribution of Brazilian deforestation fires to PM_{2.5} (Reddington et al. 2015). September marked the peak pollution month where the average person experienced PM_{2.5} concentrations of 61 µg/m³ and fires accounted for approximately 59% of the pollution mass (~36 µg/m³). In comparison to **Figure 7**, non-population weighted daily concentrations met or exceeded the US Environmental Protection Agency's 24-hour maximum standard of 150 µg/m³ on five days. Though no comparable single-day maxima

standard exists under WHO or Brazilian air quality regulations, this highlights the potential of extreme pollution exposure in low population regions closer downwind of South American fire occurrence. Furthermore, even in the absence of fires, background pollution levels are already severely degraded; the presence of fire emissions, however, significantly worsens air quality conditions. Furthermore, this analysis has focused only on the impact of local fires, yet the overall seasonality of PM matches the fire season in South America. This suggests that while local fires are enhancing exposure to pollution there is likely to be a significant contribution from longer-range fire smoke transport to the region.

To help contextualize model findings, we also examined model results for the January 2025 Los Angeles (LA) wildfire (not shown). The modelled $PM_{2.5}$ results for the LA region were muted, with a maximum population-weighted daily concentration of $29 \mu g/m^3$ on January 17. However, observational reports of the LA fire document much more extreme pollution, including a $480 \mu g/m^3$ one hour peak and a $93 \mu g/m^3$ daily mean peak on January 8th (US EPA, 2025; Briscoe and Rainey, 2025). This discrepancy likely stems from insufficient spatial and temporal resolution in both the model and the analysis region, which cannot capture the rapid and highly localized plume behaviour typical of urban or wildland-urban interface fires. It illustrates why high-resolution modelling that captures community scale air quality analysis of short-lived extreme events are needed for comprehensive impact assessments of fires as they encroach into populated regions. Benchmarking model performance against documented local maxima could guide improvements and enhance reliability for future health risk evaluations in all burning environments.



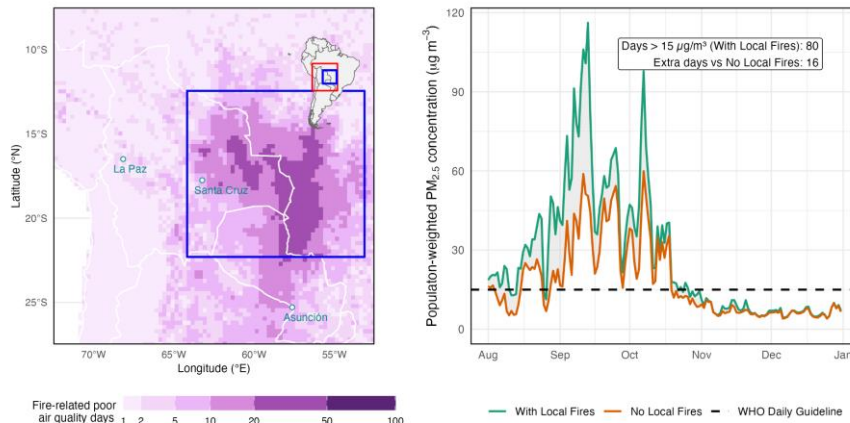


Figure 7: Poor air quality days caused by anomalous fire activity in the Pantanal-Chiquitano region during the 2024-25 fire season. **(Left panel)** shows the additional number of days exceeding the World Health Organisation (WHO) daily standard of 15 µg/m³ as a result of fire emissions occurring within the defined regions (red outlines), over and above the number of poor air quality days caused by all other sources of PM_{2.5} (e.g. industrial, transport, and residential) and from fires occurring outside of the defined regions. **(Right panel)** shows a daily time series of population-weighted PM_{2.5} concentrations (µg/m³) under scenarios that include or exclude local fires. The WHO daily standard of 15 µg/m³ is shown and days exceeding that threshold are counted as poor air quality days.

4. Diagnosing Causes and Assessing Predictability

4.1. Methods

4.1.1. Predictability of Focal Events of the 2024-25 Fire Season

4.1.1.1. Short to Medium Range Forecasts

We evaluated the capacity of two distinct methods to predict fire occurrence over short to medium -range time periods (1 to 15 days): the Fire Weather Index (FWI; van Wagner, 1987; Vitolo et al., 2020) and the Probability of Fire (PoF; McNorton et al., 2024).

FWI is a well-established empirical indicator of fire danger mostly reflecting the influence of meteorological conditions on landscape flammability. It is based on 4 prerequisite weather variables and it describes the impact that atmospheric conditions have on fuel dryness (see Section 2.1.4.1). It was originally calibrated for the boreal forests of Canada and assumes constant fuel characteristics. Due to its ease of implementation it is now widely used around the world (Bedia et al., 2015; Di Giuseppe et al., 2016; Abatzoglou et al., 2018; Vitolo et al., 2020; Jones et al., 2022). FWI is described in Section 2.1.4.1. Here, we used weather inputs from the ECMWF Integrated Forecasting system in its operational configurations at 9 km resolutions.

PoF is one of the outputs of the ECMWF Sparky fire model and aims to improve upon the fire forecasting skill of the FWI (McNorton et al., 2024; Di Giuseppe et al., 2025). The Sparky-PoF is a data-driven fire prediction system that advances on fire danger metrics by modelling not only the effect of meteorological variables on fire likelihood but also (i) the temporal evolution

of fuel load and fuel moisture content and (ii) ignition events informed by lightning forecasts, human population density, and road networks. PoF is an example of a new generation of indicators based on machine learning methods that have recently been created to produce more informative operational predictions of wildfire (Shmuel et al., 2025; Di Giuseppe, 2023). One of the practical advantages of PoF is that it can directly output a prediction of the number of fire hotspots when averaged over vast areas which is directly comparable to active fire observations. While these approaches are relatively new, they hold great promise for improving fire forecasting, particularly in fuel-limited biomes where FWI is a weaker predictor of fire activity (Bedia et al., 2015; Jones et al. 2022). ~~Sparky-PoF leverages medium-range (up to 15 days horizon) weather forecasts and fuel variables that are available from an experimental configuration where IFSI is coupled with the Sparky fire model to drive a data-driven classifier, trained on observed hotspots using from several sensors (McNorton et al. 2024) and uses~~ a XGBoost methodology (Shmuel and Heifetz, 2025; Jain et al., 2020). Predictions of PoF from Sparky showed better skills than FWI in recent events and are available operationally with forecasts up to 10 days in advance (Di Giuseppe et al., 2025).

In general, FWI is effective at capturing the immediate emergence of fire-conducive weather conditions across much of the globe. However, it does not consider the fuel build-up and the state of vegetation in specific biomes other than boreal forests, which is often a critical factor in fire occurrence. As a result, FWI-based systems may predict fire-prone conditions too far in advance of actual fire emergence, particularly in ecosystems where vegetation availability (i.e., fuel) governs ignition potential. In contrast, data-driven models like PoF, which incorporate information on both dead and live fuel moisture content are better able to reflect the delayed response of ecosystems to dry conditions. These models provide a more realistic representation of fire potential in fuel-limited landscapes or in regions where the hydroclimatic cascade delays fire onset. This is especially relevant for wetland biomes, which have been a key focus of analysis this year.

4.1.1.2. Subseasonal to Seasonal Forecast

4.1.1.2.1. Fire weather

The prediction of fire weather over sub-seasonal to seasonal (up to 6 months ahead) is a relatively unexplored field of research (Roads et al., 2005). Until recently, only a few studies had specifically examined the prediction and predictability of fire weather-related quantities and their connection to actual fire activity globally (Di Giuseppe et al., 2024). Here, we evaluate the ability of cutting-edge seasonal prediction systems to predict anomalies in the FWI, using data available through the Copernicus Emergency Management Service which uses ECMWF's SEAS5 seasonal forecasts as forcing (Di Giuseppe et al., 2024). We probabilistically quantify the likelihood of FWI values exceeding the seasonal mean prediction time steps ranging from 1 to 3 months considering a climate that spans the period 1991-2016. These predictions are not designed to inform on the exact location of fire outbreaks, but rather to serve as an indicator of landscape preconditioning to burn. The predictions highlight regions where anomalous fire weather may emerge and thus merit closer monitoring, offering an early signal of where fires could become a concern.

On seasonal timescales, patterns of fire weather are significantly influenced by large-scale climate modes such as the El Niño-Southern Oscillation (ENSO) through variation in temperature and rainfall patterns across the tropics (Latif et al., 1998; Chen et al., 2017; Bedia et al., 2018). In some tropical countries, forecasts of ENSO have been used directly to predict risk of fire and to implement preemptive fire management actions including bans on fire (Pan et al., 2018). For example, major fire anomalies and regional haze events in southeast Asia are thought to have been avoided during the 2023-2024 El Niño, following the implementation of new predictive systems and policy interventions since earlier El Niño years (e.g. 2015) (World Resources Institute, 2016). The effect of other large-scale climate modes is also

present in other world regions, such as in the case of the Indian Ocean Dipole (IOD) in the case of Australia (Harris and Lucas, 2019) and several Atlantic and Pacific oscillations in the case of Amazonia (Aragão et al., 2018). The ECMWF's SEAS5 forecasts have been shown to accurately predict the meteorological variability associated with ENSO and their effects on fire activity over timescales of 1 to 2 months ahead (Johnson et al., 2019; Di Giuseppe et al., 2024).

4.1.1.2.2. Burned Area

~~While FWI forecasts can successfully identify regions with elevated fire danger aligning with observed BA anomalies, they tend to indicate broad areas at risk and lack the specificity needed to pinpoint where fires are most likely to occur. This reflects a key limitation: translating~~ fire weather anomalies into accurate predictions of seasonal fire activity is not straightforward, as it requires incorporating additional drivers, namely fuel availability, ignition sources, and suppression capacity. Modeling the complex dynamics among fire and its bioclimatic and human drivers remains a challenge and is the focus of extensive research (e.g. Jones et al., 2022). Nevertheless, when considering forecasting ability in the long range and accuracy, climate remains the most reliable parameter among the drivers of fire activity. Accordingly, we examine the potential of machine learning techniques to forecast BA anomalies, which are being developed to provide targeted forecasts that guide the deployment and coordination of limited firefighting resources amidst increasingly synchronous wildfires (Torres-Vázquez et al., 2025a; Abatzoglou et al. 2021). We employ the model developed by Torres-Vázquez et al. (2025b), which is a hybrid approach combining dynamical seasonal drought forecasts with a statistical climate-fire model based on the Random Forest (RF) algorithm. This model leverages the Standardised Precipitation Index (SPI), aggregated over periods of 3, 6, or 12 months, to capture both antecedent and concurrent climatic conditions that influence fire activity. Calibrated with historical BA and SPI data, the RF model forecasts BA anomalies one month ahead of the fire season. The system has shown promising predictive skill, successfully capturing BA anomalies across the globe (Torres-Vázquez et al., 2025b).

4.1.1.3. Uncertainties and forecast skills

Uncertainty is a key factor in prediction and is likely to increase with forecast horizon. The forecast uncertainty is provided as the spread across a set of ensemble simulations from possible scenarios or by expressing the forecast as probability. Variability across the ensemble of forecast realizations was previously estimated to be in the range of 10%-15% for FWI (Vitolo et al., 2020), and in this study is reported as variance in the forecast values. PoF is a measure that is probabilistic in nature and is reported as probability of occurrence. For long-range predictions, uncertainty is also explicitly incorporated by expressing forecasts in probabilistic terms, specifically as the probability of exceeding (or falling below) certain thresholds, such as the upper and lower tercile.

The quality of fire forecasts is assessed by visually examining how well the forecasts capture the likelihood of key focal fire events. This approach mirrors the way fire management agencies typically interpret and use these indicators during the fire season. It is designed to partially reflect the operational context in which such indices are applied. Similarly, the seasonal predictions of FWI and the probability of above-median BA aim to demonstrate the type of information currently available to support informed decision-making for resource planning at extended lead times.

4.1.2. Identifying Causes of Focal Events

We assess the main or concurrent causes of the 2024-25 focal fire events using two complementary modelling frameworks: the Probability of Fire as part of the Sparky modelling complex (McNorton et al., 2024) and the ConFLAME attribution framework (Kelley et al., 2021;

Barbosa et al. 2025b). PoF is applied to satellite observations of active fires (Giglio et al., 2018; regridded to 0.1°) and targets a prediction of absolute fire counts on daily timescales. Meanwhile, ConFLAME is applied to satellite observations of BA from MODIS (Giglio et al., 2018; regridded to 0.5°), enabling causality analysis of fire events to key environmental and human-related causes. The ConFLAME analysis is performed on absolute BA fraction and anomalies from the 2002-2025 climatological mean and includes full regional summaries to provide broader context and to better support interpretation of region-wide drivers and trends. Used together, as in this report, the two systems provide complementary analyses of the causes of both active fire hotspots and BA anomalies.

Each model groups predictors into broader categories of causation: weather, fuel and ignitions (**Table S1 in Supplementary Material S4**). Some predictors are shared or overlap between categories due to their interconnected nature (e.g., fuel moisture and weather), but the models are designed to avoid double-counting. To identify the main causes of the fire event, PoF uses an ensemble-based gradient-boosted decision tree classifier (XGBoost), with attribution provided through SHapley Additive exPlanations (SHAP) method taken from the SHAP library (Lundberg and Lee, 2017) values to quantify the influence of each driver group on predicted [fire hotspots fires](#).

ConFLAME, in contrast, uses a probabilistic Bayesian approach to assess the contribution of each driver group to observed BA, accounting for model uncertainty and fire stochasticity. While PoF is trained globally, ConFLAME is trained separately for each region to capture regional variation in the relationship between fire drivers and BA. [Regional influence is particularly relevant for explaining the final BA as it depends on the local variations of fuel and ignition. Local ecology](#) This design is particularly important because global parameterisations, such as the use of population density as a proxy for human influence, are known to mask [regional differences in ignition practices, land-use regimes, and fire management \(Perkins et al. 2024\)](#). By focusing on regional training, ConFLAME can more directly capture how local ecological conditions (e.g. vegetation type, biomass structure). Population density in particular is split between urban and rural population densities, and BAs response to them is represented with flexible, non-linear response curves, allowing them to act as both ignition-related drivers (alongside lightning, crop and pasture fractions) and suppression/fragmentation controls (alongside crop, pasture, and urban extent). [This formulation captures region-specific, potentially humped relationships between population density and burned area, reflecting how ignition pressure and suppression capacity vary across different human-landscape contexts.](#) y shapes how vegetation and biomass affect burning (Lehmann et al. 2014) and human control can result in promoting fire (e.g. through deforestation or water extraction) or suppressing it (Andela et al. 2017). In ConFLAME causes are combined through logistic functions, with results expressed in terms of likelihoods for a detectable BA to be associated to a specific cause.

Both systems include uncertainty estimates. [PoF reflects uncertainty via probabilistic ensemble outputs and a measure is provided by the error in the predicted number of hotspots.](#) ConFLAME directly quantifies uncertainty from both drivers and model structure, providing confidence intervals for predictions. While neither system is free of limitations, this dual-model setup allows for a more robust assessment of fire causes across different spatial and temporal scales, with prediction of hotspots providing a fine-scale measure of fire activity and BA an integrated assessment of landscape impacts. A representation of PoF attribution uncertainty is made by applying a relative error derived from the comparison of observations and forecast. In ConFLAME, uncertainty has two main components. First, the uncertainty in driver-response relationships: because ConFLAME uses a maximum entropy, Bayesian inference framework, the strength and form of associations between predictors (e.g. climate, land use, population density) and burned area are not fixed but estimated from the data. This generates posterior distributions for each fitted function, which translate into confidence intervals for their contribution to fire probability. Second, the model explicitly incorporates stochastic uncertainty,

which is particularly important for extreme fire events. This reflects the inherent randomness in ignition and spread under similar conditions, and ensures that variability in event BA is not smoothed away. Together, these components produce confidence intervals that account both for parameter uncertainty and for the stochasticity of fire occurrence which is irreducible locally (i.e. gridcell), but can reduce over scale (i.e. over focal regions). They also form the basis for formal model benchmarking in **Supplementary Section 9.1**

It is important to stress that the representation of human influence on fire is very crude in both systems, with PoF relying only on static maps of population density, road networks, and land use and ConFLAME using time varying maps of urban and rural population density and land use but still missing human agency. In reality, human influence on fire is far more complex and often shaped by cultural and political constraints, which here are reflected only in observed fire occurrence rather than being explicitly represented as drivers leading to fire outcomes. These limitations are likely to introduce uncertainties that are much larger than those currently associated with the predictions, but they remain difficult to quantify due to the limited availability of detailed datasets on direct fire use, including suppression efforts. Neither system is free of limitations. Still this dual-model setup allows for a more robust assessment of fire causes across different spatial and temporal scales, with prediction of occurred fires providing a fine-scale measure of fire activity and BA an integrated assessment of landscape impacts. ConFLAME's Bayesian framework additionally helps address some of these limitations by propagating uncertainty in the estimated contribution of drivers into the modelled probability distribution, which we use directly in our analysis. This means that part of the variability in how human influences shape burning is reflected in the likelihood ranges reported by ConFLAME. Therefore, uncertainties in human-driven effects are incorporated into the analysis. However these reflect uncertainty within the model's current driver set, rather than fully capturing structural uncertainties linked to missing human agency information (e.g. suppression activities, cultural fire use, or policy change). As such, while the posterior ranges provide a useful quantitative benchmark, they likely underestimate the true uncertainty associated with human influence on fire.

The PoF from the Sparky model does not assume that each factor always pushes fire activity in the same direction. For example, while increased fuel moisture generally reduces fire activity by dampening ignition and spread, in some regions, higher antecedent rainfall can lead to greater vegetation growth, increasing available fuel and potentially resulting in more intense fires later. In fuel-limited regions, where grasses and herbaceous plants dominate, high rainfall can boost fuel growth and lead to more burning. But in fuel-rich areas with lots of trees, that same rainfall mostly increases fuel moisture, potentially decreasing fire activity. In contrast, ConFLAME allows you to specify the expected direction of influence. When a factor can both increase or decrease fire activity depending on context, those effects are represented separately in the model.

Both PoF and ConFLAME use Sparky-PoF fuel information, which inherently reflects long-term conditions, for example antecedent weather and multi-year processes are expressed in the fuel state on the day of the event. In such cases, e.g. where prior weather manifests through its influence on fuel accumulation, it is therefore categorised as a fuel driver rather than as weather itself. We assign past conditions that build up fuel loads to the fuel category, while shorter-term processes such as drying are attributed to weather, though the boundary between these two timescales is not always clear. See **Supplementary Material for Section 4** for a detailed description.

4.2. Results

4.2.1. Predictability of Focal Events

4.2.1.1. Short to Medium Range Forecasts

4.2.1.1.1. Northeast Amazonia

Between January and March, satellites detected over 30,200 fire hotspots, marking the highest number recorded for that period since monitoring began in 1999 (Eschenbacher, 2024). These fires were intensified by persistent extreme drought conditions associated with the El Niño phenomenon, which led to higher temperatures and reduced rainfall (NASA Earth Observatory, 2024c; **Figure 8**). This part of the region lies in the Northern Hemisphere tropics, where the peak of the fire season aligns with boreal winter months. The region is lesser-studied than parts of southern hemisphere Amazonia (Brando et al., 2020; Alencar et al., 2015).

Both FWI and PoF systems identified two main fire seasons in 2024: February-April and August-November. However, around 80% of the total BA concentrated in the early months of the year and only 20% during the second dry season. The total probabilities of PoF values over the focal region translates into a number of predicted hotspots and this is directly comparable to the detections from MODIS. In March, when approximately 60% of the annual BA was recorded, the PoF reached its peak predicting nearly 700 fire hotspots in a single day, closely matching the ~600 observed hotspots. While the FWI also indicated anomalously high fire risk, its peak occurred later in September. The onset of the first fire wave aligned closely with the emergence of fire-prone weather conditions (**Figure 8**), highlighting the role of weather in enabling fire activity. In this region, where over 90% of ignitions are human-driven and fuel availability remains high, atmospheric conditions primarily act as the trigger for widespread burning.

Interestingly, both the PoF and FWI systems failed to capture a lull in fire activity during the second emergence ~~in August-November~~ of fire-conducive conditions ~~showing between August and November, highlighting~~ the limitations of forecasting fire activity rather than fire danger. ~~In this region, ignitions are believed to be largely driven by escaped pasture burning Cano-Crespo et al. (2015), which typically occurs between August and October Jakimow et al. (2018). The models may have learned and reproduced this seasonal behaviour, but such patterns can be disrupted by changes in human practices.~~ One possible explanation is that these conditions fell outside ~~typical the usual~~ burning cycles, ~~—~~for example, in agricultural areas where fires are often timed around harvest, ~~the prolonged drought may have reduced crop yields and therefore fire use.~~ This ~~raises an important possibility suggests~~ that the models ~~failed to represent missed~~ the quiet September period because they ~~have incorporate~~ only limited information on human ignition patterns, land ownership and ~~land~~ use types, and ~~lesser-studied less documented~~ factors such as fire suppression, policy interventions, and ~~managed or~~ cultural burning practices, ~~underseering Lapola et al. (2023). These gaps underscore~~ the need for improved ~~datasets on human activity—data that, which~~ could significantly ~~improve enhance~~ fire prediction (Jones et al., 2022).

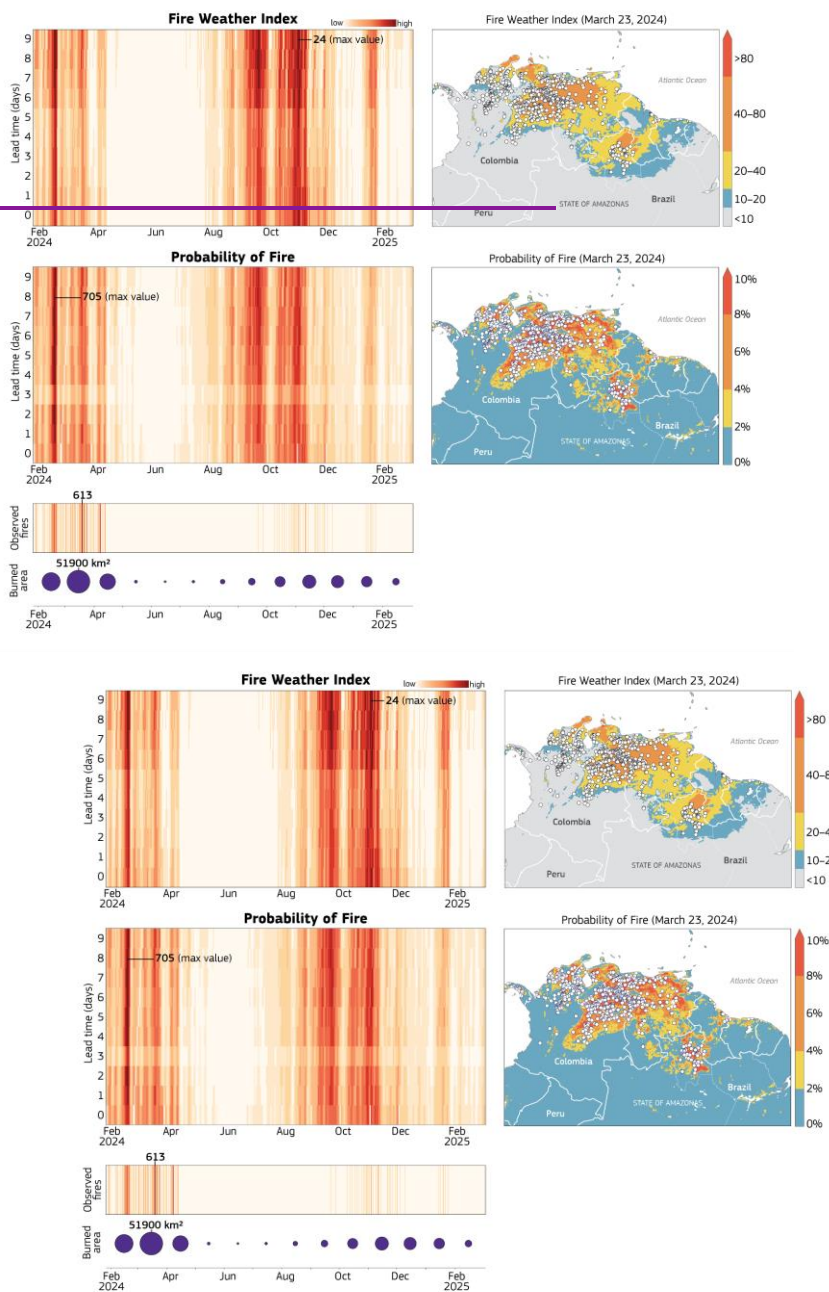


Figure 8: Northeast Amazonia forecasts of the FWI and PoF with lead times up to 10 days prior for the period February 2024-February 2025 as an average value over the focal area.

The total percentage of PoF values over the focal region translates into a number of predicted hotspots and this is directly comparable to the detections observed by MODIS. The x-axis corresponds to specific dates throughout the year, while the y-axis denotes either observations or the time leading up to the date when a forecast was generated. The vertical colour coherence allows for quick identification of the time windows of predictability associated with the observed fire activity both provided in terms of number of detected active fires per day and total monthly BA (circles). The maps represent a snapshot in time at day 0 to allow the comparison of the spatial distribution of the forecasts and the recorded fire activity by MODIS.

4.2.1.1.2. *Pantanal and Chiquitano*

The Pantanal and Chiquitano have been enduring a prolonged dry period since 2019 leading to the 2024 worst water [availability](#) crisis ever recorded in the biome (World Wildlife Fund, 2024). Notably, the Pantanal did not experience its typical flood season in early 2024 and the average area covered by water during the first four months was smaller than that of the previous year's dry periods (Van Dijk et al., 2025). By the end of May 2024, almost the entire Pantanal and Chiquitano region was classified as experiencing extreme drought, the second-highest classification of drought intensity on the Integrated Drought Index (NASA Earth Observatory, 2025a, 2024c). As the Pantanal is a wetland ecosystem, the establishment of dry conditions is a prerequisite for the onset of fire activity. A full hydrological cascade must occur before widespread burning can take place: prolonged precipitation deficits must lead to the reduction of flooded areas, their replacement by grasslands, and the progressive desiccation of both live and dead vegetation. This sequence introduces a natural delay, which explains why fire activity in the region peaked in August and September, well after the onset of dry weather in June (**Figure S11 in Supplementary Material S4**).

The total percentage of PoF values over the focal region translates into a number of predicted [hotspots/cells with fires](#) and this is directly comparable to the detections from MODIS. The most severe PoF forecast, predicting 971 [hotspots/fire cells](#), closely matches the 885 observed in late August. At large scales, the FWI offers a useful overview of fire-conducive weather conditions. However, it is the inclusion of fuel characteristics in the PoF that provides the finer spatial granularity (maps in **Figure S11 in Supplementary Material S4**) needed for more accurate and actionable fire risk assessments.

4.2.1.1.3. *Southern California*

California is arguably one of the most extensively studied regions in terms of shifts in fire regimes (see, e.g., Billmire et al., 2014; Littell et al., 2016; Williams et al., 2019; Swain et al., 2025). In 2024, Southern California experienced severe burning in September, with a total of 1,200 km² burned. Although these fires fell within the typical fire season, the total BA was unremarkable for the region compared to previous years. However, the most significant fire event took place much later, in January 2025, well outside the typical fire period, when the Palisades and Eaton fires broke out in Los Angeles county. The events sparked widespread public debate about how prepared we are to anticipate off-season fires (Woolcott, 2025).

As shown in **Figure S12 (Supplementary Material S4)**, fire-prone weather conditions persist across much of the year, extending well into autumn, a reflection of the expanding fire season driven by climate warming. Yet, in regions like Southern California, fire prediction based solely on weather indicators is often inadequate. The primary causes of the severity of these events was an intensification of the hydrological cycle that exacerbated both wet and dry extremes. Southern California experienced an unusually wet antecedent period prior to intense drying in an unusually dry winter, which created an accumulation of dry fuel setting the ideal conditions for intense fire activity (Swain et al., 2025). Fuel accumulation is a persistent feature throughout the fire season, and therefore does not result in a large difference between the PoF and FWI forecasts when averaged over the Mediterranean areas of California. However,

its inclusion in the prediction system allows for the identification of zones with higher susceptibility, which are clearly visible in the accompanying map. Neither the FWI nor PoF metrics could provide adequate warning regarding the magnitude of the winter fire event affecting the wildland-urban interface. These events were driven by atmospheric phenomena influenced by steep orography, which are not resolved by current weather forecasting models. The lack of the required resolution impacts equally on empirical and machine learning methods. This highlights the need for improved high resolution forecasting for fire danger in the wildland-urban interface

4.2.1.1.4. Congo Basin

The 2024 dry anomaly in Central Africa has been partly attributed to the co-occurrence of a positive El Niño phase and a warm Indian Ocean Dipole (McPhaden et al., 2024). These conditions tend to shift the West African monsoon northward, leading to suppressed precipitation over the Congo Basin during the core of the rainy season, a pattern observed globally in recent years (Toreti et al., 2024). This event also aligns with a broader trend of a lengthening and intensifying dry season in the Congo rainforest. Satellite analyses over the past few decades show that the dry season is starting earlier and ending later, increasing the region's vulnerability (Jiang, 2019). There are typically two main fire seasons in the Congo Basin: from December to March north of the equator, and from June to September south of the equator. In the equatorial zone, however, fires are not naturally occurring, as precipitation is distributed throughout the year. The expansion of dry seasons both north and south of the equator has led to a situation where fire seasons in the Congo Basin now span almost the entire year with peak activities between July and August and December and March. Compounding this, a decline in lightning activity over the region (Chakraborty and Menghal, 2025) suggests that fires are increasingly of human rather than natural origin. This combination of persistent drier-than-average conditions and human-driven ignition means that fire activity is now widespread and weakly correlated with weather patterns. As a result, predictions have a very short predictability window of only a few days (horizon at correlation of lines in **Figure S13**). The detachment of fire activity from natural conditions in the Congo Basin presents a significant challenge for forecasting (**Figure S13**). In these regions, the discriminatory power between fire-prone and non-prone conditions is greatly reduced, and both FWI and PoF tend to overpredict fire occurrence. In particular, the FWI fails to capture the complex interactions among fuel availability, ignition sources, and human activity. Although the PoF incorporates some of these elements, it also struggles in regions where human presence and behaviour are rapidly changing. In such cases, while most burning occurred outside conflict zones (Trigg et al., 2012), the broader instability in the region may still influence local fire activity, challenging the predictive capabilities of data-driven systems if not trained on the most recent data. This limitation is especially ~~pronounced~~ evident in areas where natural ignitions are infrequent, and fuel dynamics, rather than weather alone, drive fire occurrence and behaviour (Figure S13). In the future, these inherent limitations may be addressed by incorporating more fire-specific socio-economic factors. We are aware that datasets providing more detailed information on human practices are becoming available (Kasoar et al., 2024), and these may help constrain and improve forecast skill going forward. Such datasets could also provide a valuable basis for further exploring the links between fire predictability and human influences, building on the data presented in this report.

4.2.1.2. Seasonal Predictability from Fire Weather Forecasts

The year 2024 has been officially declared the warmest year on record, surpassing previous temperature benchmarks (WMO, 2025; NOAA, 2025a). This exceptional warmth has been driven not only by long-term global warming (IPCC, 2023), but also by a combination of short-term ocean-atmosphere anomalies. In particular, extensive and persistent oceanic heat waves have been observed across multiple ocean basins, contributing to elevated sea surface temperatures (Holbrook et al., 2019). These marine heatwaves have been further reinforced

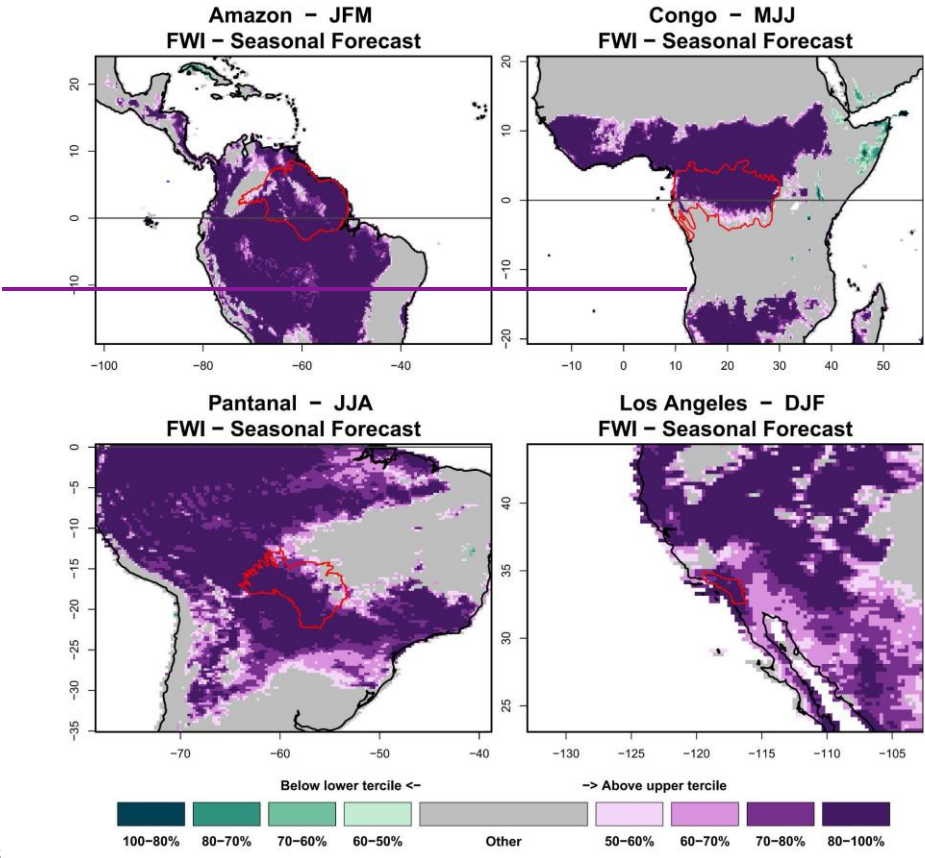
2237 by an unusual reduction in low-level cloud cover over parts of the Atlantic Ocean, allowing for
2238 increased solar radiation absorption at the ocean surface and amplifying the warming (Ceppi
2239 and Nowack, 2021).

2240 Given this overall picture, seasonal forecasts of FWI anomalies successfully captured the
2241 broad regional patterns of elevated fire danger, particularly in Northeast Amazonia and parts
2242 of Bolivia and Venezuela (**Figure 9**). These forecasts aligned with the widespread drought
2243 and above-average temperatures linked to the strongest El Niño since 2015, a concurrent
2244 positive Indian Ocean Dipole, and record-breaking ocean heatwaves. Together, these factors
2245 intensified drying across equatorial South America, expanding fire-prone conditions well
2246 beyond the regions that ultimately experienced the most extreme burning.

2247 All forecasts issued one month before the fire season showed high confidence (between 60
2248 and 90%) in the development of above-normal conditions in our focal regions, all exceeding
2249 the 66th percentile of climatological values.

2250 **Figure 9** demonstrates both the strengths and limitations of FWI-based seasonal forecasts.
2251 While they provide valuable early warnings by detecting fire weather anomalies, their broad-
2252 scale nature can lead to overestimations of fire impact if not combined with information on fire
2253 susceptibility. These findings reinforce the value of FWI in anticipating periods of increased
2254 landscape flammability, but also highlight the need to more appropriately model anomalies in
2255 fuel load and moisture and to integrate non-climatic factors, such as ignition sources, land use
2256 practices, suppression capacity, and landscape accessibility, into fire impact forecasting
2257 models to improve precision and operational relevance. Future seasonal-scale forecasts may
2258 seek to implement PoF as a predictive tool, which improves upon FWI by tracking fuel loads
2259 and moisture and thus the legacy effects of antecedent conditions on landscape flammability.

2260 **Figure S14** presents an example of the burned-area anomaly forecasting system using our
2261 hybrid dynamical and Random Forest (RF) approach (**Section 4.1.1.2.2** and Torres-Vázquez
2262 et al., 2025b). The maps illustrate the predicted probability of a BA anomaly and whether these
2263 predictions could trigger alerts for BA anomalous seasons within a potential early-warning
2264 system. Following Torres-Vázquez et al. (2025b), alerts are issued when predicted
2265 probabilities exceed thresholds optimized to balance correct detections and false alarms. For
2266 the 2024 season, anomalies in South America, notably in drought-affected regions influenced
2267 by El Niño conditions, were reasonably well anticipated. However, in other regions, particularly
2268 parts of Africa including the Congo basin, there were numerous false alarms, reflecting current
2269 limitations in fully capturing regional complexities and non-climatic fire drivers. This first
2270 implementation demonstrates operational potential, and future refinements (such as
2271 incorporating extended fire records and adjusting region-specific thresholds) could enhance
2272 skills by reducing false positives.



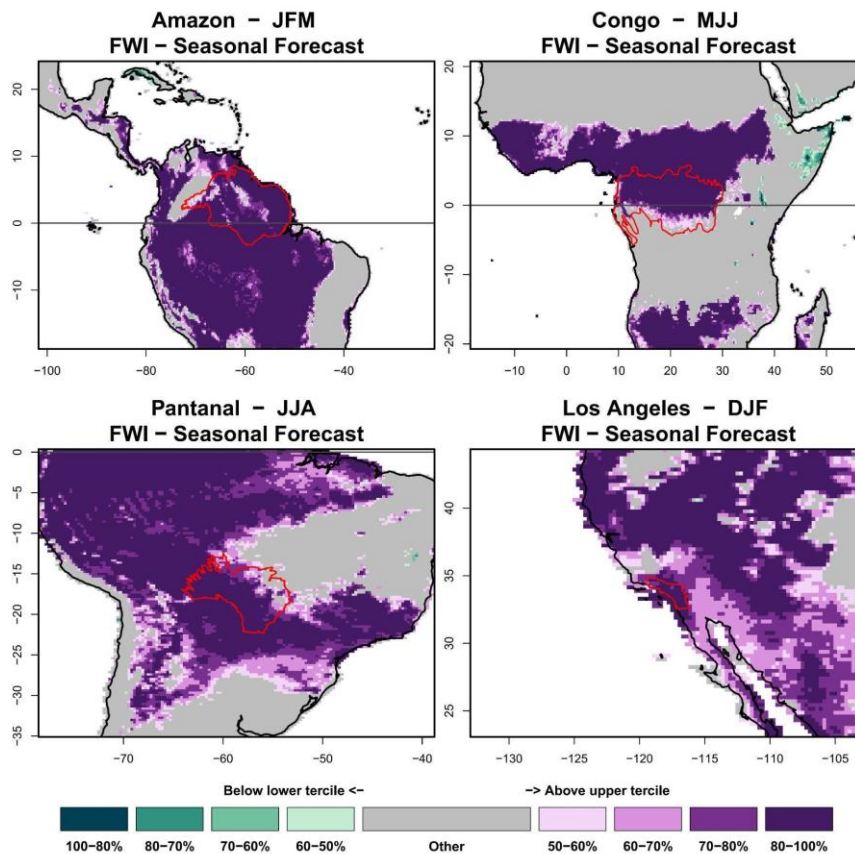


Figure 9: Seasonal prediction of Fire Weather Index (FWI) during the periods relevant to our focal events, presented in probabilistic terms that indicate the likelihood of an increased or decreased anomalous fire season.

4.2.2. Identifying Causes of Focal Events

Weather, fuel, and ignitions are the three primary controls factors influencing the occurrence and intensity of fire events (Di Giuseppe et al., 2025). These broad categories can be further examined to pinpoint individual factors, for example, precipitation and temperature within the weather category or fuel moisture from dead and live vegetation in the fuel category. Analysing the single factors can give an idea not only of the probability of the fire to occur but also on their intensity and behaviour. For example anomalies above the expected climate (here 2003-2023) in the moisture of dead fuel, due to its lower moisture content and higher combustibility, often plays a significant role in determining ignition potential. Low live fuel moisture increases vegetation flammability, thereby contributing significantly to greater fire severity and intensity.

Beyond this descriptive approach, the PoF and ConFLAME causality models enable a probabilistic attribution of fire occurrence to the three primary controls. These models provide attribution even when no fire is recorded: a low probability across all controls reflects an accurate forecast of no fire, while a high probability without observed fire activity could point

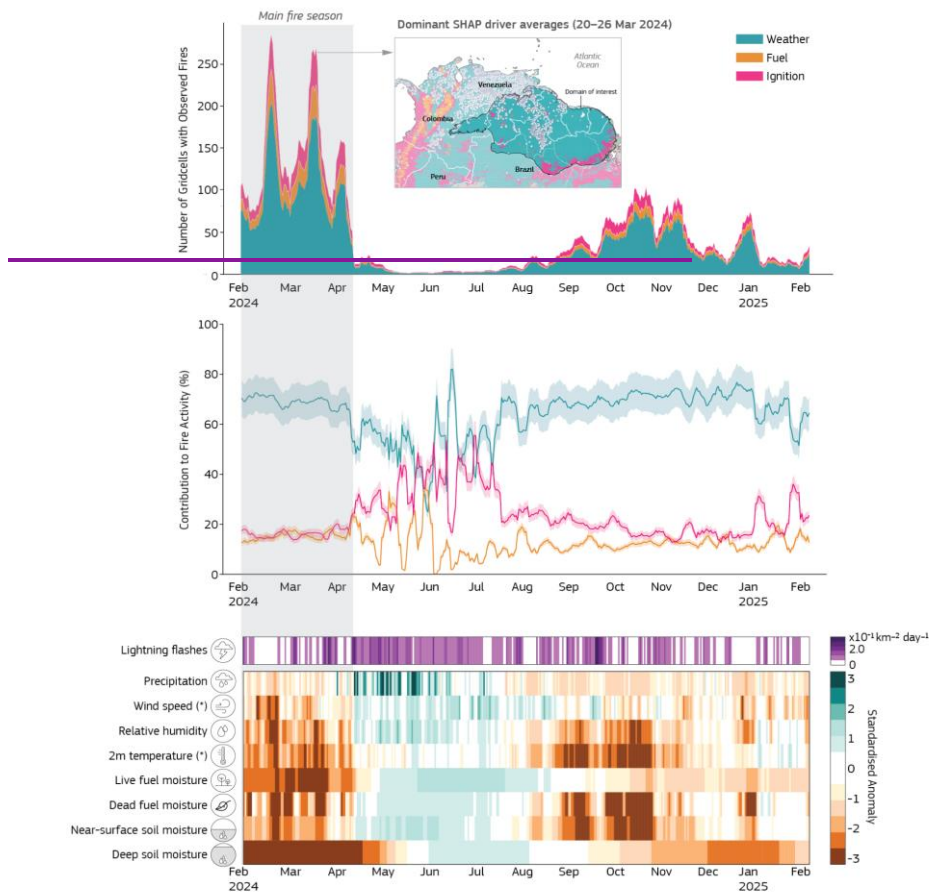
to successful suppression efforts, fire-prevention policies, or other unaccounted human factors not included in the model forecasts. The discrepancies between the model prediction and the observed fire activity (cells with identified fire hotspot or BA anomalies) are included to provide a measure of the model uncertainties.

4.2.2.1. Causes of Fire Hotspots during Fires during Focal Events

4.2.2.1.1. Northeast Amazonia

According to our Sparky-PoF analysis, the extreme fire activity during the 2024-25 fire season in Northeast Amazonia (described in Section 2.2.2.1), was predominantly driven by anomalous dry weather. Northeast Amazonia experienced an exceptionally severe fire season between January and April (Figure 10), driven by extreme drought which started in 2023, intensified by the combined effects of El Niño and the Atlantic Meridional Mode, which brought unusually high temperatures and suppressed rainfall. At the peak of the season, during the week of 20-26 March, nearly 2,000 fire-hotspots were observed. Fires were fueled by prolonged and intense drying across the entire landscape, which made vegetation highly flammable and enabled rapid fire spread across large areas. On the most severe week of burns our causation analysis shows that weather conditions were the dominant factor, accounting for about 60% of fire activity, while fuel availability and ignition sources each contributed around 20%. During the first part of the year the exceptional dryness meant that soil humidity levels and moisture in both dead and live vegetation fell to among the driest 2% of historical conditions, while deep soil moisture dropped below 1%. The time series of lightning activity (Figure 10, bottom panel) further illustrates that ignitions in the region are predominantly human-driven. During the May-August period, lightning activity is high and is linked to storms and rainfall, which tend to suppress fire ignition and spread. As a result, even though lightning increases the relative contribution of ignition to predicted fire activity, doubling its weight to around 40%, this is not reflected in actual fire occurrence or BA. A second, less intense onset of fires occurred between September and January. This was driven by a more superficial drying of the landscape that did not extend into deeper soil layers. Unlike the earlier season, which was associated with hydrological drought, this later period was more reflective of meteorological drought (precipitation deficit).

Northeast Amazon



Northeast Amazon

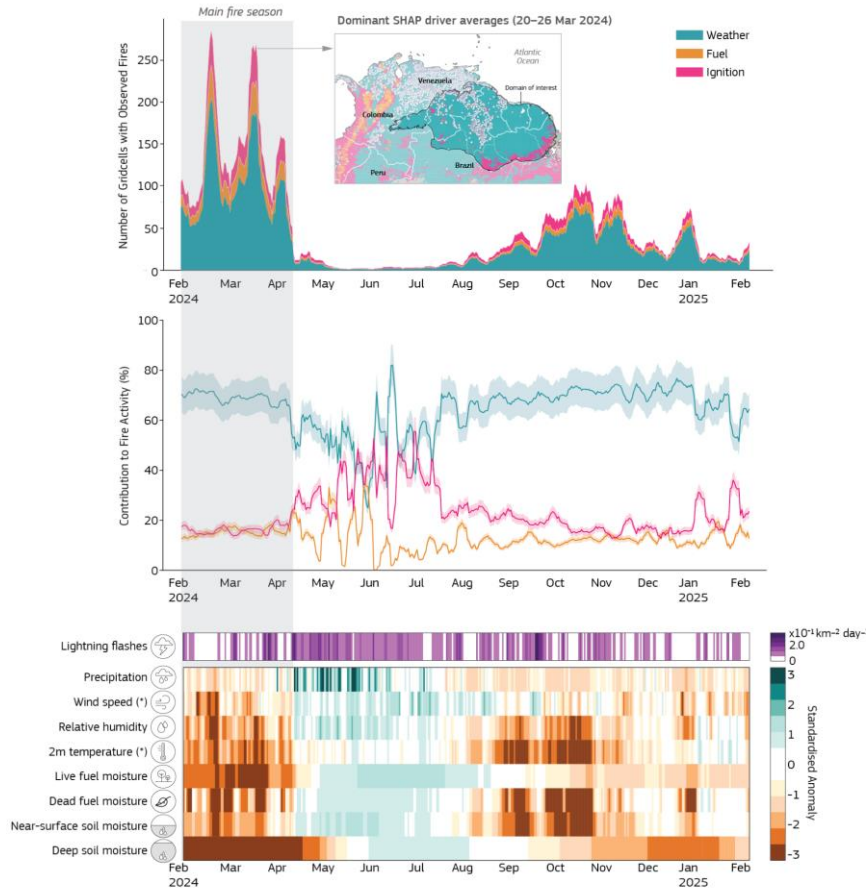


Figure 10: Drivers explaining fire prediction in Northeast Amazonia. Daily fire activity and contributing drivers from February 2024 to February 2025. **(Top panel)** Daily count of grid cells with detected fire hotspots, stacked by dominant driver category, fuel, weather, or ignition/suppression. A dominant driver is assigned only if its contribution exceeds 50% of the total attribution; otherwise, the grid cell is left unclassified (gray). An inset map shows the spatial distribution of dominant drivers during the peak fire week, highlighting regional heterogeneity in fire causation. **(Middle panel)** Relative contributions (%) of each driver category to predicted fire occurrence, with shaded bands indicating model-observation uncertainty. **(Bottom panel)** Standardized anomalies (in units of standard deviation) for each input driver variable, including lightning flash density. Asterisks (*) indicate reversed anomalies.

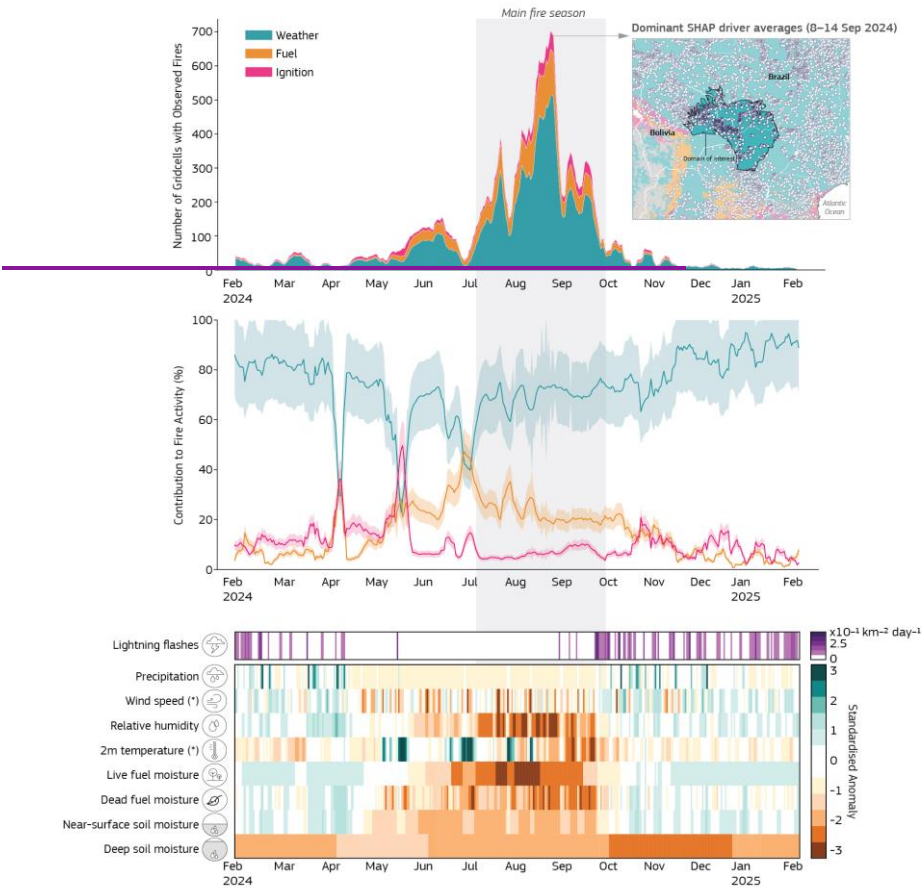
2337 4.2.2.1.2. Pantanal and Chiquitano

2338 According to our Sparky-PoF analysis, the extreme fire activity during the 2024-25 fire season
2339 in the Pantanal-Chiquitano (described in **Section 2.2.2.2**), was mainly the result of extremely
2340 dry weather which had started since 2023. Drought conditions affecting the Pantanal and
2341 Chiquitano continued into the early months of the 2024-25 fire season following multiple years
2342 of below-average rainfall (**Figure 11**). Although the year began with relatively moist surface
2343 conditions, deep soil moisture remained in the driest 15% of observed records or 1-2 standard
2344 deviations below the mean (**Figure 11**). A wet phase in February-April allowed moisture
2345 transfer from the atmosphere to surface fuels, but it did not infiltrate deeply into the soil. As a
2346 result, when surface conditions dried out again at the beginning of June, vegetation quickly
2347 became highly flammable and primed to ignite.

2348 While fire activity in this region was predominantly controlled by weather (71% mean
2349 contribution throughout the year), the role of fuel became increasingly important during the
2350 most intense burning phases (up to 40% during the most intense week between 5 and 14
2351 August 2024). In fact, the contribution of fuel conditions doubles during these peak events,
2352 indicating that the persistence of fire-conducive weather over time, rather than the specific
2353 daily weather, plays a dominant role in driving the most severe fires.

2354 In the Pantanal and Chiquitano, the lack of correlation between fire occurrence and natural
2355 ignition sources, such as lightning density (**Figure 11**, bottom panel), is even more evident
2356 than in other regions. When lightning does occur, it is typically accompanied by rainfall due to
2357 the convective nature of tropical storm systems, further reducing the likelihood of fire ignition.
2358 The only notable 'dry lightning' event, observed in mid-May, caused a spike in the modelled
2359 PoF which translated into a spike of fire activity that was observable though small in
2360 magnitude. Humans are the main source of ignitions in the region (Menezes et al., 2022) and,
2361 while weather remains the main driver of fire activity overall, fuel conditions are playing an
2362 increasingly important role in determining the severity and extent of extreme fire events
2363 (**Figure 11**).

Pantanal and Chiquitano



Pantanal and Chiquitano

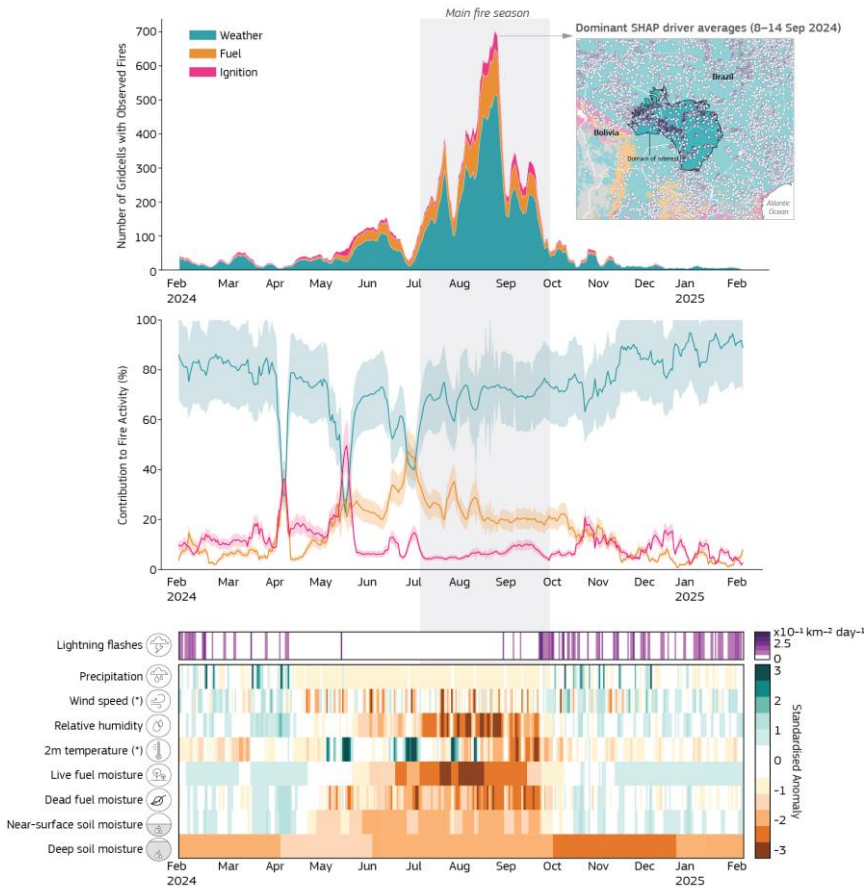


Figure 11: Drivers explaining fires (model cells with identified hotspots) in the Pantanal-Chiquitano (as for Figure 10).

4.2.2.1.3. Southern California

According to our Sparky PoF analysis of the extreme fire activity during the 2024-25 fire season in Southern California (described in Section 2.2.2.3), the results point to a combination of drivers, weather, fuel, and ignitions, each playing an almost equal role in creating the fire prone conditions observed during the two major events in January 2025 (Palisades and Eton fires).

Early in the 2024-25 fire season, Southern California was emerging from a two-year period of very wet conditions, with deep soil moisture levels at 2 to 3 standard deviations wetter than the climatological average (McNorton et al. 2025)(Figure S15). During the summer of 2024, lightning may have contributed to ignitions, although in these areas most fires are typically

human-induced. Overall, fire activity remained relatively low and below the climatological average.

However, the Palisades and Eaton fires in January 2025 were well outside the typical fire season. These fires were clear outliers in terms of their seasonality, triggered by a rare alignment of short-lived but intense fire-prone conditions while fuel moistures remained low (Figure S15). Between 5 and 25 January, favourable weather, fuel availability, and ignition sources aligned leading to create ideal conditions for ignition and rapid fire spread.

In the week preceding the fires, fire weather conditions contributed around 40% to the predicted fire probability, fuel availability 30%, and ignition sources the remaining 20%. Despite the generally moist deep soil conditions, a brief but extreme episode of surface drying (reaching 3 standard deviations below normal) combined with unusually strong winds (also 3 standard deviations above average), was sufficient to create highly flammable conditions at the wildland-urban interface, enabling the fires to ignite and spread rapidly.

4.2.2.1.4. Congo Basin

According to our Sparky-PoF analysis, the extreme fire activity during the 2024-25 fire season in the Congo basin (described in Section 2.2.2.4), was the result of the extreme drought that has affected the regions in recent years.

In 2024-2025, fire activity in the Congo occurred year-round in a region marked by abundant and widespread vegetation cover. The spring wet season (March-May) did not materialise due to extreme and persistent drought conditions. As a result, the second wet season later in the year also brought limited relief, leaving deep soil layers significantly dry (up to 2 standard deviations below climatological norms). The region remains in a prolonged state of water deficit until now (Figure S16).

Throughout the year, weather conditions were the dominant and most stable factor influencing both the number, intensity and duration of fire events. A combination of low rainfall (67% below the climatological average) and elevated temperatures (90% above the climatological average) led to sustained drying of both vegetation and soil, placing them among the driest 2% and 1% of the climatological record, respectively. These conditions maintained highly flammable landscapes across the region (Figure S16)

Most fire ignitions in the Congo basin can be attributed to human activity. Although lightning occurs year-round (Figure S16), it is more frequent during the wet season due to the convective nature of tropical precipitation. However, during these wetter periods, high moisture levels typically prevent fire ignition and spread. In contrast, during prolonged dry spells, even a small number of human-caused ignitions can trigger widespread and persistent fire outbreaks, owing to the highly combustible state of the vegetation.

4.2.2.2. Causes of Burned Area Anomalies during Focal Events

4.2.2.2.1. Northeast Amazonia

According to our ConFLAME analysis of the extreme BA during the 2024-25 fire season in Northeast Amazonia (described in Section 2.2.2.1), weather conditions explained about 40-60% of the BA anomalies, though with fuel conditions acting as an important determinant cause during the periods with greatest fire extent (Figure 12). In the peak month of March 2024, BA exceeded the long-term average (2002-2024) by over 12,000 thousand km². Nearly half of the March 2024 anomaly could be attributed to fuel conditions, while weather anomalies potentially accounted for between 50% and 150% of the BA anomaly (a high-end value of 150% would suggest that weather alone would have caused anomalies exceeding the

2424 observed values, but below-average ignitions moderated the BA response; **Figure 12**). During
2425 the secondary peak in BA anomalies during October-November, fuel and weather contributed
2426 similarly with fuel rising in importance due to the insufficient water recharge from the wet
2427 season. Weather and fuel together accounted for between 1,000 thousand km² and over
2428 10,000 thousand km² of BA anomalies.

2429 Consistent with active fire analysis (**Section 4.2.1.1**), fuel played a key role in determining the
2430 geographical distribution of BA during the 2024-25 fire season (**Figure S19**). This is visible in
2431 northern regions such as the Forest-Savanna Transition Zone in northern Venezuela and
2432 southern Guyana, and the Northern Amazonia Savannas of Roraima and northern Pará,
2433 where savanna outcrops are surrounded by rainforest (see maps in **Figure S19**). In the forest
2434 landscapes, fuel anomalies and fire weather anomalies drove the predicted anomalies in BA.
2435 Interestingly, predicted BA anomalies were large in some parts of the region (e.g. Suriname)
2436 but went undetected by the MODIS BA product. The causality framework is very confident in
2437 its prediction, raising the question of whether detections were missed, possibly due to dense
2438 canopy and persistent cloud cover (Giglio et al., 2006).

2439 Despite widespread BA in early 2024, many parts of the region remained largely unburned.
2440 Understanding why is as important as knowing what drove the fires. Our analysis shows that
2441 in areas with very low BA fraction (less than 0.5% of burnable area), no single factor (fuel,
2442 weather, or human activity) clearly limited fire spread (refer to **Figure S18**). Instead, a
2443 combination of factors, such as low ignition rates, patchy fuels, or short dry spells, likely
2444 prevented fires from taking hold. On the other hand, in the most severely burned areas (top
2445 5% of BA), the relative importance of fuel and weather was reversed compared to broader
2446 patterns. Here, fuel moisture emerged as the primary driver of BA. Drier conditions could have
2447 increased BA by 30-40%. Weather still played a role contributing an additional 20% increase,
2448 but its influence was secondary to that of fuel.

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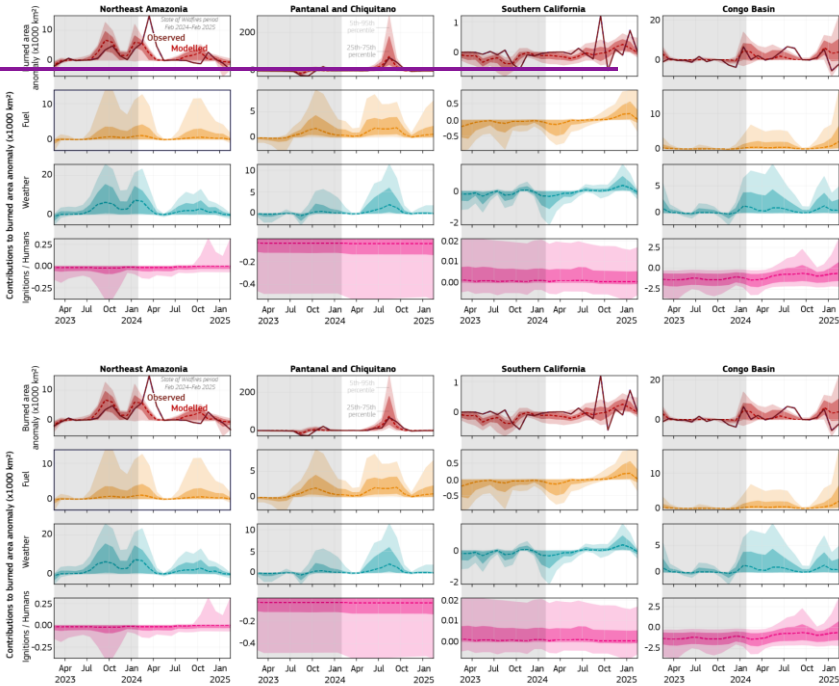


Figure 12: Anomalies in burned area (BA) and driver contributions for each focus region during 2024, relative to the 2002-2024 average. Columns represent regions; rows show different variables. **(Top row)** Modelled and observed BA anomalies, expressed thousand km². Model output shows median, interquartile range (shaded), and 5th-95th percentile range (lighter shading). **(Other rows)** Modelled contributions to BA anomalies from fuel conditions, fire weather, and human/ignition-related factors, also shown in thousand km². These panels highlight which drivers contributed most to regional fire deviations from the historical average in 2024.

4.2.2.2.2. Pantanal and Chiquitano

According to our ConFLAME analysis of the extreme BA during the 2024-25 fire season in the Pantanal-Chiquitano (described in **Section 2.2.2.2**), weather conditions explained half of the BA anomalies and fuel conditions explained almost 30% (**Figure 12**). June, July, and August accounted for the most extensive burning in the Pantanal, with 25-75% of the landscape experiencing some fire activity, even if large parts featured only small anomalies. The peak occurred in June, when the BA exceeded climatological values by more than 5,000 thousand km² (almost triple than the annual mean). This anomaly was primarily driven by weather conditions (50-60%) with fuel (10-20%) and ignition (10-20%) contributing equally. Although fire weather remained favourable in September and October due to persistently high temperatures, overall fire activity was lower than during the earlier peak (**Figure 12 and S20**).

We found that ignition sources contributed only 10-20% to the anomalously high fire activity in 2024. However, we caution that our modelling framework only partially captures ignition

dynamics, particularly those related to human activities such as farming. This limited representation is reflected in the wide uncertainty range assigned to ignition within the causality framework. Key factors like land clearing, water extraction, and the proximity of ignitions to protected areas are known contributors to extreme fires in the Pantanal (Barbosa et al., 2022) and they are not fully accounted for in our analysis.

Regional differences in fire drivers were evident (**Figure S20**). Fuel conditions played a key role in the fine-scale geographical distribution of BA anomalies. Exceptionally dry fuels affected the Chiquitano dry forests in the east, while weather was the dominant driver in upland regions along the edge of the Pantanal wetlands, such as the Serra do Amolar hills in western Brazil. The most extreme fires were observed where these two influences overlapped, where vegetation was both unusually flammable and atmospheric conditions were conducive to burning.

As for what prevented the fires from becoming even more severe (**Figure S18**), no single factor alone limited fire spread, even in the areas that burned most intensely. However, small shifts in conditions, such as drier weather, drier fuels, or fewer land-use barriers, could have led to 2-12% more BA in the model cells experiencing the greatest fire anomalies regions (top 5% of anomalies).

4.2.2.2.3. *Southern California*

According to our ConFLAME analysis of the extreme BA during the 2024-25 fire season in Southern California (described in **Section 2.2.2.3**), the most important cause of the extent of burned areas was fuel (30% to 60%) closely followed by weather (20-40%), while ignitions (20%) was less pronounced than in previous years and acted as reducing factor (**Figure 12**). During January 2025, unusually dry fuel conditions played a key role in promoting BA anomalies, explaining up to 500 km² of the 800 km² of the anomalous BA in that month. Fire weather conditions, starting as early as October 2024, were also anomalous versus previous years. Focusing on the areas with the most extensive burning (top 5% of BA), we found that anomalies could have been 30-60% larger under drier fuel conditions and more extreme fire weather, with an additional 5% increase if fuel availability anomalies had also been higher (**Figure S17 and Figure 4.11**). The substantial suppression efforts deployed is unaccounted for in our modelling framework and could be one of the possible reasons the fires did not escalate even further.

4.2.2.2.4. *Congo Basin*

According to our ConFLAME analysis of the extreme BA during the 2024-25 fire season in the Congo basin (described in **Section 2.2.2.4**), weather conditions explained about 30-60% of the BA anomalies, with fuel conditions acting as an important secondary control during the periods with greatest fire extent (**Figure 12**). Fuel conditions in the Congo Basin remained relatively stable throughout 2024, contributing between 10-35% to fire activity year-round. In contrast, the influence of weather conditions varied more substantially, with virtually no fire-conducive weather in October-November, outside the typical fire season, and moderate levels (5-15%) during peak fire periods, particularly in January and July (see also **Figure S17**). July stood out as the month with the largest deviation from typical fire patterns. During this time, fuel conditions and fire weather contributed almost equally to the BA (**Figure S22**).

ConFLAME indicates widespread anomalous BA across the southern part of the Congo basin. These model estimates of BA are larger than the BA detected by satellites (**Figure S22**). Dense canopy in these remote regions may have led to missed detections of BA. Particularly high fire-conducive conditions were predicted across much of southern Democratic Republic of the Congo (DRC), as well as northern Angola and parts of the Republic of the Congo. However, two notable pockets, in the far northeast of the basin, around the border of northeast DRC and

2522 South Sudan, and a smaller zone just east of the border between the DRC and Republic of
2523 the Congo in the north, did not emerge in our analysis.

2524 Despite these broad areas of fire-favourable conditions, fires did not become much larger in
2525 many places. The key reasons for this were moisture and weather limitations. Looking at areas
2526 at the top 5% of burning, up to 15% more fire could have occurred if fuel had been even drier
2527 or if atmospheric conditions had been slightly more favourable.

5. Attribution to Global Change Factors

Many of the direct drivers and controls on fire events, outlined in **Section 4** (e.g. weather, fuel, moisture, ignition and suppression), are influenced by global change factors such as climate and land-use change. Since the pre-industrial era, global mean temperature has increased by ~1.3°C (Betts et al., 2023; Forster et al., 2025), with greater rates of warming at higher latitudes, adding potential for fuel drying. Climate change has also resulted in altered precipitation patterns, with total rainfall and dry season length increasing or decreasing variably across regions (Polade et al., 2014; Swain et al., 2018; IPCC, 2023a). Meanwhile, changes to fuel load and ignition rates are driven by emissions, climate change and land-use change, with varying effects regionally (Foley et al., 2005; Finney et al., 2018; Roms, 2019; Wang et al., 2024).

5.1. Methods

5.1.1. Overview of Attribution Approaches

Fire is a complex phenomenon that impacts societies and ecosystems in many ways, from the extent of BA to the severity of individual fire events. Different user groups seek information on different aspects of fire risk, whether policymakers, communities, fire managers, litigators, or those working to build a broader scientific evidence base. To provide results relevant for a wide range of stakeholders, we apply various modelling approaches to fire attribution, drawing on different metrics and attribution techniques, to build a more comprehensive understanding of human influence on recent extreme fire activity. Our approach includes analyses of fire weather indices and BA, alongside a range of attribution metrics suited to these different contexts. Our BA attribution also provides the evidence, in the form of a calibrated probabilistic model, needed to perform future risk projections in **Section 6**.

While most attribution research has focused on the contribution of anthropogenic climate change, humans influence fire occurrence and risk in multiple other ways: the direct influence of people via activities such as land use change and landscape configuration; changes in ignition probability, fire suppression, among others. Considering human-driven climate change separately to changes in human activity, in addition to their combined effect, allows us to disentangle the contributions of local [socio-economic](#) and global environmental change.

[A key challenge that has rarely been addressed before is how to represent socio-economic influences on fire within a framework that can also attribute climate change \(Huntingford et al. 2025\). This is a relatively new area of research, and progress has been limited by data availability, the complexity of human–fire interactions and how to represent these in models. The absence of variables capturing human agency, such as deliberate ignitions, suppression, or governance, has made it difficult to capture the ways people influence fire \(Perkins et al. 2024\). For instance, previous attempts to incorporate socio-economic drivers have often relied on global parameterisation of population density as a simple predictor of fire activity. However, this relationship is shaped by local cultural, political, and economic contexts and can therefore give misleading results. In this report we extend last year’s framework by taking a more detailed and regionally grounded view of how people influence burning through using urban and rural population densities \(see **Section 5.1.3**\) and using non-linear response mechanisms \(**Section 4.1.2**\). This represents an important step towards capturing socio-economic influences more realistically, but it is still only a partial representation. Considerable work remains to incorporate human agency and broader socio-economic context \(Millington et al. 2022\), and we see this as an active area of development for future reports.](#)

Understanding the influence of people or climate on fire and its drivers is inherently challenging, given the complexity of fire processes and the interactions between natural and human systems. Integrating these range of complementary methods - each with its own strengths and limitations, additionally helps build confidence in attribution results that no single method could provide alone. We can therefore identify where there is broad agreement across methods. [This dual focus on climate forcing and socio-economic drivers also allows us to identify where there is broad agreement across methods and to highlight where gaps remain, particularly in representing local socioeconomic influences on burning.](#)

To quantify the different ways people affect fire, we apply four types of attribution in this report (**Table 4**), designed to meet diverse user needs and to align with the modelling frameworks currently available:

- i) Firstly, our attribution to *anthropogenic climate forcing* explicitly targets the changes driven by anthropogenic greenhouse gas emissions and land-use change, following the IPCC WGI definition (Hegerl et al., 2009; Mengel et al., 2021). We prescribe these emissions in a model to specifically isolate human forcing from natural variability (**Section 5.1.2 and 5.1.3**).
- ii) Our attribution to *total climate forcing* considers changes driven by climate change since the pre-industrial period, including both anthropogenic climate forcing and natural variability in line with the IPCC WGII and the Intersectoral Impacts Model Intercomparison Project 3a (ISIMIP3a) definition of climate change impact attribution (IPCC, 2023b; IPCC 2023c; Mengel et al., 2021). This involves comparing simulations driven with historical reanalysis, our factual, to a detrended counterfactual simulation, where the trend in each climate variable is removed (with both simulations including historical transient land-use change). Therefore only the impacts of climate change are attributed, not distinguishing between anthropogenic or natural causes (Mengel et al., 2021; Burton, Lampe et al., 2024). We perform this between 2003-2019, the overlap between available counterfactual simulations and satellite data used for training in Burton, Lampe et al., 2024.
- iii) Our attribution to *socio-economic factors* is applied via the same set of simulations as our attribution to *total climate forcing*. We isolate the role of socio-economic factors by comparing the early industrial period to the late industrial period (1901-1917 versus 2003-2019) using detrended ISIMIP3a data, in which only land-use and population density are allowed to change (Burton, Lampe et al., 2024). [with one of our frameworks incorporating a more detailed representation of local population-fire interactions. This definition, used consistently in both Burton & Lampe et al. \(2024\) and Jones et al. \(2024\), provides a transparent and comparable first-order approach to including direct human pressures in attribution analyses.](#)
At the same time, this should be viewed as a step in an active but still incomplete field of research. Full socio-economic attribution remains an emerging area, and further work is needed to incorporate processes such as ignition practices, suppression, and governance, as well as to capture how human-fire relationships evolve over time. Our framework therefore represents a step forward, internally consistent, peer-reviewed, and comparable across regions, while highlighting where future developments will be most impactful.
- iv) Our attribution to *all forcings* compares the early industrial period in the counterfactual scenario to the last industrial period in the factual scenario, which gives the net effect of all forcings combined (*anthropogenic climate forcing + total climate forcing + socio-economic factors*).

The attribution methods described above enable us to assess the influence of climate and socio-economic forcings on fire in each focal region with respect to three different target variables:

- i) **Extremes in fire weather during 2024-25.** ~~The FWI is a weather-based indicator of landscape flammability and can provide insight into how fire-prone conditions are likely to be affected by a changing climate.~~

Using the HadGEM3-A large ensemble, we attribute changes in the probability of extreme fire weather conditions to *anthropogenic climate forcing*. This analysis specifically targets the months identified as extreme for each focal event as outlined in **Section 2.2.2** focusing on sub-regional extremes that occur in the model grid cells with the highest FWI values (top 5% of all regional grid cells). By focusing exclusively on these areas of most severe fire weather, this approach provides a proxy for understanding how each forcing influences the locations and times of highest fire risk within the region. We used this methodology as in last year's report. See **Section 5.1.2** for details.

- ii) **Region-wide extreme BA during 2024-25 focal events.** Event specific BA reflects how climate and human factors jointly influence the actual extent of burning during major fire events, offering a direct measure of fire impact on people and ecosystems.

Using the ConFLAME model framework we attribute changes in the likelihood of the 2024-25 observed total BA across the entire focal region to *anthropogenic climate forcing*, *total climate forcing*, *socio-economic factors*, and *all forcings* combined. Like our FWI analysis focuses on the observed peak burning months and captures the overall influence of each forcing on the extent of fire activity at the regional scale. See **Section 5.1.3** for details.

- iii) **Background changes in BA this century** using median monthly over recent decades. Background BA shows how climate change is reshaping regional fire regimes over the long term, revealing gradual shifts in baseline fire activity that may go unnoticed in year-to-year variability.

Using fire-enabled dynamic global vegetation models (DGVMs) participating in the Fire Model Intercomparison Project (FireMIP), we attribute changes in median monthly BA averaged over recent decades (2003-2019) to *total climate forcing*, *socio-economic factors*, and *all forcings* combined. This approach provides context on longer-term background fire activity and applies the same methodology as last year's report, though this year focussing on specific focal regions. See **Section 5.1.4** for details.

In each approach we include an explicit estimate of uncertainty. We use bootstrapping to give uncertainty estimates for the FWI Risk Ratios (RR) defined as the ratio between the probability of seeing the observed FWI with the target forcing vs without anthropogenic climate forcing, reported here at 90% confidence intervals. ConFLAME is designed as an uncertainty quantification model (as per our driver assessment, **Section 4.2.4**), giving the likelihood of all possible BA outcomes for each region based on a probabilistic analysis of past burn patterns and environmental conditions. We combine the information from the FireMIP models in a weighted multi-model ensemble to give uncertainty ranges across the models. Each result therefore presents a 5-95th percentile probability estimate.

For consistency with last year's report we also report attribution estimates based on methods used in the State of Wildfires 2023-24 report (Jones et al., 2024b):

- iv) **Sub-regional extreme BA during 2024-25.** We attribute changes in the likelihood of extreme BA occurring within the model grid cells with the highest BA (top 5% of all regional grid cells), focusing on areas where fire activity was most spatially concentrated during peak burning months. This analysis uses the same ConFLAME simulations and forcing scenarios as the region-wide BA attribution and provides insight into how forcings affect the most severely impacted locations within the region. See **Supplementary Text S5.2.3** for discussion of results.

In the coming years, our project seeks to incorporate attribution results based on a broader set of Earth System Models (ESM) to better sample the structural uncertainty arising from differences in process representation across different models (i.e. beyond HadGEM3-A). In this report, we introduce results based on the one ESM as follows:

- v) **Background changes in fire weather this decade.** Using the Canadian Earth System Model (CanESM5; Swart et al., 2019), we attribute changes in the frequency of extreme fire weather to *total climate forcing* with the [Canadian Fire Weather Index \(FWI\)](#), identifying how the likelihood of extreme fire weather has changed by comparing the frequency of high [Fire Weather Index \(FWI\)](#) values in pre-industrial and present-day climates. Our analysis covers the years 2016 to 2025, focusing on the climatological months of peak burning during the 2024-2025 fire season. See **Supplementary Text S5.1.2** for methodology and **Supplementary Text S5.2.2** for discussion of results.

Table 4: Summary of the attribution approaches used in this report. See **Table S2** for a breakdown on the what each attribution type includes and what each modelling targets.

Term	Definition	Experiments compared	Framework	Application
Event attribution for fire weather and burned area				
Anthropogenic climate forcing	Change in fire weather driven by anthropogenic emissions from greenhouse gases, land-use change and aerosols. As per (Ciavarella et al., 2018; Li et al., 2021)	Factual: HadGEM3-A_ALL: with natural forcing plus human emissions Counterfactual: HadGEM3-A_NAT with natural-only forcing from solar variability and volcanoes	HadGEM3-A attribution ensemble. 0.83 x 0.56 degree resolution	Fire-weather (FWI)
			ConFLAME (Kelley et al. 2021; Barbosa et al. 2025b) with merged ERA5/HadGEM3-A product	Burned Area with ConFLAME
Impacts attribution for fire weather and burned area				
Total climate forcing	Changes in FWI since pre-industrial	Factual (2016-2025): present-day climate from CanESM5 SSP585 Counterfactual 1850-1859): Pre-industrial climate simulation	CanESM5 CMIP6 ensemble	FWI
	Changes in BA due to climate change, irrespective of the cause of warming. As per ISIMIP (Intersectoral Impacts Model Intercomparison Project) (Mengel et al., 2021 and Frieler et al., 2024)	Factual (2003-2019): present-day climate (driven by GSWP3-W5E5 reanalysis), CO2, land-use and population Counterfactual (2003-2019): Historical climate detrended using seasonally-varying regression on global mean temperature (ATTRICI method, CO2 fixed at 1901 value, present-day land-use and population	ISIMIP3a impact attribution. 0.5 degree resolution	FireMIP ensemble and ConFLAME
Socio-economic factors	Changes in BA due to land-use and population change. As per (Burton, Lampe et al., 2024)	Counterfactual (1901-1917): Warming trend removed using ATTRICI method, fixed 1901 CO2, limited land use and population change Counterfactual (2003-2019): Warming trend removed using ATTRICI method, fixed 1901 CO2, present-day land use and population		
All forcings	Changes in BA due to climate, land-use and population change. As per (Burton, Lampe et al., 2024)	Counterfactual (1901-1917): Warming trend removed using ATTRICI method, fixed 1901 CO2, limited land use and population change Factual (2003-2019): Historical climate driven by reanalysis	ISIMIP3a impact attribution	FireMIP ensemble

5.1.2. Attributing Extremes in Fire Weather during 2024-25

We use two complementary approaches to attribute changes in the probability of high fire weather, measured usingby the Canadian Fire Weather Index (FWI)₁ to anthropogenic climate change. The first method uses a targeted large-ensemble weather model simulation to assess the influence of climate change on the 2024/25 fire seasons directly. The second method applies a longer-term, probabilistic framework using simulations from a fully coupled Earth system model.

2714
2715 The first approach follows the same methodology used in the previous State of Wildfires report
2716 (Jones et al. 2024b). This is an established approach to attribute changes in the probability of
2717 high fire weather, measured using FWI, to anthropogenic climate forcing. This method has
2718 been previously used by the World Weather Attribution [Initiative](#) (Barnes et al., 2023; Barnes
2719 et al., 2024; Barnes et al., 2025), using outputs from the HadGEM3-A large ensemble
2720 (Ciavarella et al., 2018). Our approach builds on the methodology introduced by Stott et al.
2721 (2004) for attributing extreme weather events, and it has been applied in other attribution
2722 studies targeting fire weather, such as Li et al. (2021).

2723
2724 As outlined in **Section 4.1.1**, the FWI is used operationally and in research contexts to rate
2725 fire danger based on meteorological conditions. Due to the availability of model output, [which](#)
2726 [is typically only available on a daily temporal resolution](#) variables we use maximum daily
2727 temperature at 1.5 m as a proxy for noon values, total daily precipitation, mean daily relative
2728 humidity at 1.5 m, and mean daily wind speed at 10 m, following Perry et al. (2022). We
2729 calculate the daily FWI for the months of 2024-25 peak BA anomaly for each focus region,
2730 using the same month and region for validation over the historical time series (1960-2013).
2731 Note that at time of writing, data for HadGEM3-A was only available till the end of 2024, so we
2732 do not report on Southern California fires using this method.

2733
2734 We validate and bias-adjust the model estimates of high FWI for the period 1960-2013 by
2735 comparing a 15-member HadGEM3-A ensemble with ERA5 reanalysis data (C3S, 2024)
2736 representing "observed" FWI. The 0.25 degree resolution observed FWI from ERA5 was
2737 coarsened by linear interpolation (calculated by extending the gradient of the closest two
2738 points) to match the 0.5 degree model grid. We compare the time series of individual
2739 components of the FWI ([Figure S49-S55S52-S59](#)), and the distribution of the modelled and
2740 observed FWI ([Figure S56-S58-S67-S70-](#)), and apply a simple linear regression to find the
2741 bias correction required for the 2023 model output. Before bias-adjustment, the modelled FWI
2742 is generally higher than the observed FWI for Amazonia and Congo, which modelled FWI
2743 compares more favourably to ERA5 in the Pantanal. The correction adjusts the trend and
2744 absolute value while maintaining variability, and the model successfully reproduces the
2745 observed distribution after applying the correction in each region (see **Supplementary Text**
2746 **S9**).

2747
2748 For the events occurring in the 2024 fire season, we calculate the FWI from the HadGEM3-A
2749 model simulations comprising 2 experiments of 525 members each, one driven by all forcings
2750 including historical greenhouse gas emissions, aerosols, zonal-mean ozone concentrations,
2751 land-use change and natural forcing (ALL), and a second counterfactual simulation with
2752 natural-only forcing from solar variability and volcanic emissions, and 1850 land-use (NAT)
2753 (see **Table 4**). By applying the bias-adjustment from the previous step, and comparing the fire
2754 weather in the two simulations to the 2024-25 observed FWI from ERA5, we calculate the
2755 change in probability of high fire weather due to anthropogenic climate forcing. The standard
2756 definition of "high fire weather" that we use is the 95th percentile of [daily Fire Weather Index](#)
2757 [\(FWI\) daily FWI](#) values across all grid cells and days during the season. However, as in last
2758 year's report and in Burton et al. (2025), when the region is small or when climate conditions
2759 significantly influence the higher FWI in our counterfactual, leading to few ensemble members
2760 reaching higher FWI values, we need to adjust our definition of extreme. In this year's
2761 assessment, we apply the 90th percentile threshold for the Northeastern Amazonia and Congo
2762 regions, as the differences between the factual and counterfactual ensembles are so large
2763 that very few counterfactual members reach the 95th percentile of the factual distribution,
2764 making the calculation of risk ratios unreliable.

5.1.3. Attributing Region-wide Extreme BA during 2024-25

We use the ConFLAME framework for direct BA attribution. For this report, we apply two configurations of the ConFLAME attribution framework to attribute anomalies in BA fraction during the peak burning months of the 2024-25 fire season:

- A **near real-time (NRT) setup** for targeting anthropogenic climate forcing, which largely mirrors the configuration used in the drivers attribution section (see **Section 4**), assesses how human influences affected the likelihood of BA via meteorological driver of fire conditions observed during the specific 2024 events. This setup targets the actual environmental conditions leading up to and during the events, providing the most up-to-date picture of climate and socioeconomic influences. By focusing on the precise timing and location of the event, the NRT configuration provides an up-to-date and high-resolution picture of how anthropogenic climate forcings have influenced the likelihood of extreme fire activity.
- The **Inter-Sectoral Impact Model Intercomparison Project (ISIMIP) 3a setup**, previously used with ConFire in last year's report. This setup enables the analysis of how often fire events such as those in 2024 might occur under environmental conditions from 2002 to 2019. While 2024 itself is excluded, we look for similar events in this earlier period to understand how likely they would be without the recent changes in climate and land use. This broader, long-term setup helps us assess how the background risk is shifting over time and complements the more event-specific analysis shown earlier. This setup also directly links to the future projections presented in **Section 6**, which also use ISIMIP. As an addition to last year's report's set up, our ISIMIP set up also includes (i) changes in land use and cover (measured as the difference between tree cover and agricultural fraction since the previous year) ~~in the direct socioeconomic~~ and (ii) a separation of urban and rural population density in the socio-economic forcing attribution (see **Table S3**). The former allows us to capture direct effects of land-use change on fire, rather than only static land-use states, while the latter enables us to represent non-linear relationships between population density and fire, including both positive and negative influences. Together, these developments represent methodological advances introduced in this year's report. As noted in Section 4.1.2, ConFLAME is optimised regionally and, for population density in particular, provides a more effective parameterisation than would be possible with a globally optimised model (Perkins et al., 2024).

As each configuration uses data that is somewhat similar to our Fire Weather (in the case of NRT) or FireMIP (when using ISIMIP) set ups, neither setup is fully independent of our other two modelling approaches. However, the fire modelling in ConFLAME captures different components of fire than FWI or FireMIP by attributing BA during the events themselves. The advantage of ConFLAME is that it bridges the gap between event-focused real-time attribution and global process-based fire models. That said, future iterations would benefit from incorporating more independent, preferably observation-driven input datasets to improve robustness and reduce potential structural alignment across methods.

Each attribution experiment involved training ConFLAME using "observed" or reanalysis driving data against MODIS BA (as described in Section 4). We then ran the framework with factual driving data followed by a separate run counterfactual with the effect we aim to attribute (e.g., all forcings, climate, or socioeconomic drivers) removed. We conducted paired ConFLAME factual and counterfactual predictive model simulations at monthly resolution, using a structure similar to that in **Section 4.1.2**, with specific drivers grouped into controls in **Table S1** and evaluated the model following Barbosa (2025; **Section 4.1.2**). We separately train ConFLAME on 50% of the data between 2003-2011 and perform evaluation on years 2012-2019. Further details of the model fitting and validation can be found in **Supplementary Text S5.1.3** and **Supplementary Text S9.1.2.2**, respectively.

To determine the impact of total climate forcing, socioeconomic factors and total forcing on increased BA during our focal events using the ISIMIP configuration, we conducted paired sampling of monthly BA in the target months (see Table 4). Total climate forcing's factual driving data uses the same 2003-2019 GSWP3-W5W5 reanalysis data used for training for factual, while we use detrended data for the counterfactual, whereas socioeconomic used detrended data 2003-2019 for factual and 1901-1917 for counterfactual. Total forcing used 2003-2019 from GSWP3-W5W5 for the factual and 1901-1917 from detrended GSWP3-W5W5 for the counterfactual. We used paired sampling to account for uncertainty in the relationships between drivers and BA, ensuring co-variation between experiments (as in Kelley et al., 2021). In total, we drew 1,000 samples across the 17 years of each simulation, resulting in 17,000 paired samples. [Evaluation of performance can be found in Supplementary Text S9.3.](#)

[We use two key metrics to assess how our target factors have influenced BA during extreme fire events.](#) We report attribution [metrics/results](#) both for the entire region (reported in the main text, **Section 5.2.2**) and for "sub-regional extremes" - the grid cells with the top 5% of BA, to also assess how anthropogenic factors may have influenced the most severely affected areas (in **Supplementary Material S5.2.2**). [The \). We use three complementary metrics to assess how our target factors have influenced burned area \(BA\) during extreme fire events:](#)

1. **Amplification Factor (AF)** [tells us/quantifies](#) how much [bigger-\(larger or smaller\)-the](#) BA was because of [a specific factor/the considered driver](#). It [works by](#) [comparing/compares](#) factual BA for [BAsevents](#) as large or larger than [what was/the](#) observed [during the target months versus/with](#) counterfactual. [Observed BA is calculated in a manner consistent with the model outputs, by averaging BA across either the entire region or the top 5% of BA within the target region and month. Observations are derived from monthly MCD64A1 data. In near-real-time \(NRT\) mode, we do this for the specific year of interest. In the ISIMIP setup, we compare across many years \(2003-2019\).](#) An AF [greater than](#) 1 means [climate change increased BA. For example, indicates an increase due to the driver \(e.g. AF of 2 means → twice as much area burned, or an AF = 1/2 → half as much burned area due to the target factor\).](#) We calculate this across our model simulations and report both the central estimate (median) value and the range of uncertainties based on 10th to 90th percentiles.
2. **Likelihood of Attribution** [the probability that BA was higher under the factual \(with-forcing\) world than it would have been in the counterfactual \(without-forcing\) world, expressed as a percentage.](#)
3. For the NRT set up, we can also use the **Risk Ratio (RR)** which [expresses how many times more likely an event of comparable BA was under factual versus counterfactual conditions/.](#) [Similarly to Section 5.1.2, it compares the chance of seeing the observed BA under today's climate to the chance under a climate without human influence. A RR above 1 means climate change made the event more likely; a RR below 1 means it made it less likely.](#) ~~Because~~

[Observed BA is calculated in a manner consistent with model outputs by averaging BA either across the entire region or, for sub-regional extremes, across the top 5 % of grid cells. Observations are taken from the monthly MCD64A1 dataset. In near-real-time \(NRT\) applications, the comparison is made for the specific year of interest; in the ISIMIP setup, the comparison spans 2003–2019.](#)

~~Because~~ the Early Industrial factual simulation in our ISIMIP setup includes no human influence on the climate, we first adjusted the target event's BA to the level expected without climate change. This adjustment involved identifying the percentile of the observed BA in the factual simulation, and then finding the BA at that same percentile in the counterfactual simulation.

~~For the NRT set up, we can also use the Risk Ratio (RR), which shows how much more (or less) likely the target factor made a fire event of this size. Similarly to Section 5.1.2, it compares the chance of seeing the observed BA under today's climate to the chance under a climate without human influence. A RR above 1 means climate change made the event more likely; a RR below 1 means it made it less likely.~~

5.1.4. Attributing Background Changes in Burned Area this Century

We assess how BA has changed over recent decades due to climate and socio-economic drivers using the FireMIP ("Fire Model Intercomparison Project") attribution framework developed by Burton, Lampe et al. (2024). This method uses state-of-the-art global FireMIP models, employing each model's native fire scheme, to estimate the contribution of different drivers to BA by comparing simulated fire activity under different ISMIP3a experiments ~~setup~~. We quantify the effect of *climate forcings* on BA by comparing the present-day factual burned area to the present-day counterfactual BA. The effect of *socio-economic forcings* is assessed by comparing the present-day of the counterfactual simulations to the early-industrial of the counterfactual simulations since long-term climate is stationary in these simulations. Lastly, we find the effect of *all forcings* by comparing the present-day factual BA to the early-industrial counterfactual BA.

The attribution focuses on changes in median monthly BA during 2003-2019 and uses a weighted multi-model ensemble, where weights reflect each model's ability to reproduce observed regional fire anomalies in GFED5 and FireCCI datasets. Uncertainty is evaluated through random resampling of the weighted ensemble, including a stochastic parameter that captures uncertainty in overall ensemble performance. This weighting reduces the influence of models that fail to capture local fire–population or fire–bioclimate relationships, but does not fully resolve structural gaps in the ensemble. In particular, weak performance for socio-economic drivers may widen overall uncertainty without attributing it to the specific process. As a result, FireMIP provides a conservative estimate of regional-scale fire responses, complementing our more detailed but regionally focused approaches.

All results are reported as relative anomalies, and uncertainty is assessed via a random resampling of the weighted ensemble, including a stochastic parameter which accounts for uncertainty on the performance of the entire ensemble. This approach provides a robust and conservative estimate of trends, particularly suited to assessing regional-scale fire responses.

In contrast to last year's report, where results were reported for IPCC AR6 regions containing the focal fire zones, this year we refined the analysis by tailoring the attribution directly to the specific areas featured in the report. This regional adjustment enhances the relevance and interpretability of the attribution results for each case study.

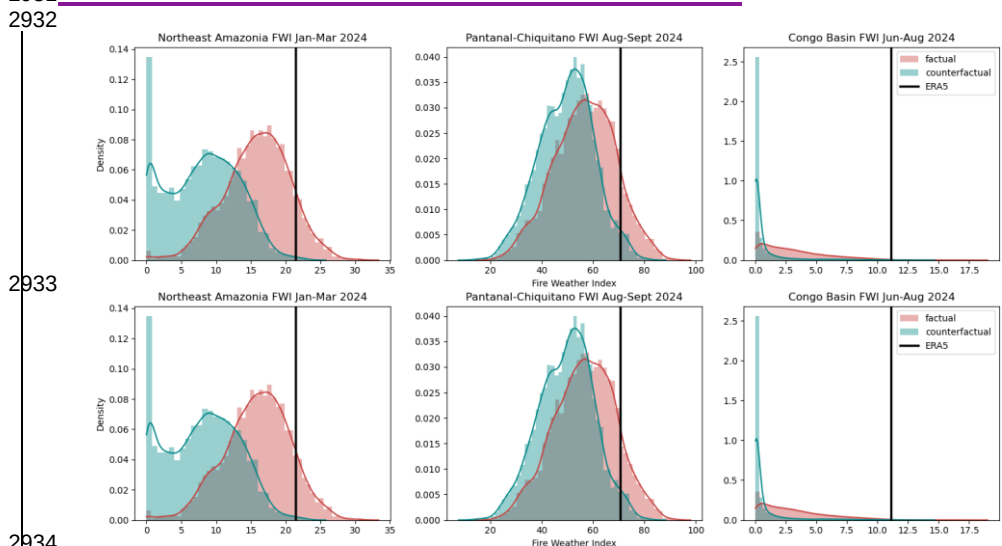
For full details on the method, model evaluation, and baseline results across all IPCC regions, see Burton, Lampe et al. (2024).

5.2. Results

2919 **5.2.1. Extremes in Fire weather during the 2024-25 Focal Events**

2921 **5.2.1.1. Northeast Amazonia**

2923 We find that the fire weather conditions in Northeast Amazonia during January-March 2024
2924 were significantly more likely due to anthropogenic climate forcing, with the probability
2925 of experiencing fire weather at or above the levels observed during the event being 32 to 73
2926 times higher more likely in the factual simulations compared to the counterfactual simulations
2927 (Figure 13). A substantially larger proportion of the factual ensemble exceeds the observed
2928 90th percentile of FWI from the ERA5 reanalysis than in the counterfactual ensemble (Figure
2929 13), indicating that high fire weather conditions during early 2024 were much more likely in a
2930 climate influenced by anthropogenic emissions.



2935 **Figure 13:** High fire weather conditions in 2024/25: Probability distributions of FWI in the
2936 HadGEM3 ensemble for the focal fire season in each region, comparing simulations with
2937 anthropogenic and natural forcings (red; factual) to natural-only forcings (teal; counterfactual).
2938 Black line shows ERA5 reanalysis. The x-axis shows the regional average of high-percentile
2939 FWI days: 89th percentile for Jan-Mar in Northeast Amazonia (left), 95th percentile for Aug-
2940 Sep in the Pantanal and Chiquitano (middle), and 90th percentile for Jun-Aug in the Congo
2941 Basin (right).

2943 **5.2.1.2. Pantanal and Chiquitano**

2945 The high fire weather conditions experienced during the peak anomaly in fire activity in August-
2946 September 2024 were 4.2-5.5 times more likely due to anthropogenic climate forcing (Figure
2947 13). While this increase in likelihood is smaller than that estimated for Northeast Amazonia,
2948 the narrower range suggests we have greater confidence that human influence increased the
2949 probability of extreme fire weather conditions in this region.

2950 Our results largely agree with the rapid attribution analysis from the World Weather Attribution
2951 (WWA) initiative (Barnes et al., 2024), though with smaller uncertainty ranges, WWA found
2952 that the accumulated fire weather conditions, represented by the June Daily Severity Rating
2953 (DSR), were 4.6 (1.1 to 20) times more likely due to human-induced climate change. The DSR,
2954

a fire-suppression oriented rescaling of the FWI, is commonly used to assess the cumulative fire weather danger over monthly timescales (Van Wagner, 1987). WWA focused on June conditions because of their role in setting up the severe fire season that followed, and their direct relevance to the large BA that severely impacted wildlife and livelihoods in the Pantanal. Observations also indicated a decrease in annual rainfall of -23.5% (-46% to +5%) in the region, though this trend was not reproduced by climate models (Barnes et al., 2024).

5.2.1.3. Southern California

Due to the lack of data availability from HadGEM3-A for 2025, we were unable to perform bespoke FWI attribution analysis for Southern California. However, in previously published analysis, the rapid attribution study by WWA (Barnes et al., 2025) found that the extreme fire weather conditions (peak FWI) in the coastal southern California ecoregion surrounding Los Angeles during January 2025 were 1.37 (0.48 to 3.6) times more likely in comparison to the pre-industrial climate, suggesting that climate change may have [leadled](#) to a moderate increase in fire weather, though causing a reduction in fire weather is also plausible and within the confidence range. As the impacts of Los Angeles fires related to extreme single days of wildfire spread, the monthly maximum FWI value averaged over the study region was used here. This result is complemented by the increasing likelihood of an extended dry season in the region. Decreased October-December precipitation allowed for protracted fuel drying, resulting in a more likely overlap between dry conditions and the winter Santa Ana winds. Observed trends (ERA5) in the October-December standardised precipitation index found that the dry conditions leading up to the LA fires were 2.4 (0.33 to 20.9) times more likely than in the pre-industrial climate. Using analogue-based attribution (Vautard et al., 2016), the cut-off-low circulation pattern associated with the strong Santa Ana winds around Los Angeles was found to have increased in likelihood by 2.5 (0.4 to 17) times.

5.2.1.4. Congo Basin

The high fire weather conditions observed across the Congo Basin during June-August 2024 were unusual in both the factual and counterfactual simulations. Our analysis indicates that these conditions were 3.0-8.0 times more likely due to anthropogenic climate forcing (**Figure 13**). The entire FWI distribution in the factual ensemble is shifted toward higher values compared to the counterfactual ensemble. This means that across the full range of fire weather conditions, the probability of conditions conducive to burning is substantially greater in a climate influenced by human emissions.

5.2.2. Region-wide extreme BA during 2024-25

5.2.2.1. Northeast Amazonia

We find strong evidence that anthropogenic climate forcing contributed to increased regional BA during the January-March 2024 fire season in Northeast Amazonia. Our analysis shows a 96% likelihood (very likely under IPCC definitions of confidence) that BA was higher than it would have been without anthropogenic climate forcing (**Figure 14**). We estimate that regional BA was approximately 4.3 times larger (our *Amplification Factor - how much larger or smaller burned area is because of climate change*) than it would have been in a counterfactual world without anthropogenic climate forcing (**Figure 14; Table 5**), with a 90% confidence range of 1.02 to 25.32. While the central estimate suggests a quadrupling of BA, the wide uncertainty range reflects the natural variability of fire processes. Nonetheless, even the lower bound supports a small but clear increase.

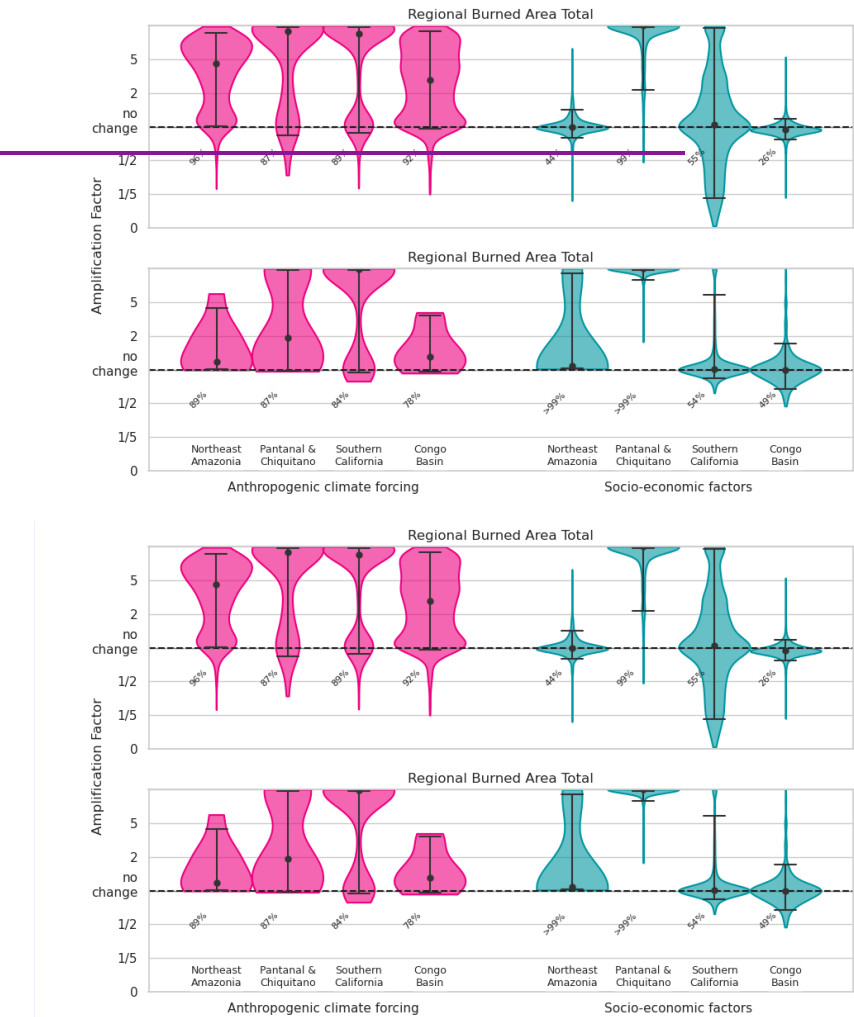
We assess the risk ratio, the likelihood of an event like January-March 2024 occurring under current climate conditions versus a pre-industrial baseline (**Table 5**). Based on historical data provided as evidence for the model, we estimate that a similar event is now 2.1 times more

likely due to anthropogenic climate forcing. This figure captures the longer-term climate signal that would shape the overall frequency of such events. When we control for meteorological variability by comparing simulations with and without anthropogenic forcing but using identical weather patterns from 2024, we see slightly stronger effects (**Table 5**). The risk ratio rises to 2.7, and the upper bound of our Amplification Factor increases dramatically (over 100-fold in some ensemble members). This suggests that climate forcing alone could account for much, or possibly all, of the burning under certain conditions, although the central estimate remains close to our previous assessment.

Climate influence was widespread across Northeast Amazonia, most of the entire region showing a greater likelihood of increased BA due to anthropogenic forcing (**Figure 15**). The strongest attribution signal occurred in the Southern Guiana Shield Fringe Forests, where climate change was very likely ($\geq 90\%$ confidence) to have increased BA. These forests are particularly important due to their extensive areas of primary rainforest and high ecological sensitivity. In contrast, attribution confidence tapered to around 70-80% in the Guiana Coastal Plain, and only a few localized areas, particularly in savanna mosaics, showed weak or no signal.

~~The region's ecological heterogeneity, encompassing floodplain forests, natural grasslands, and savanna formations, means fire impacts vary considerably. Some savanna systems are naturally adapted to low intensity surface fires (Alvarado et al., 2020; Pivello et al., 2021), but increased frequency and intensity of burning can overwhelm their resilience. Fire sensitive ecosystems, such as humid forests and wetlands, are even more vulnerable, with increased fire pressure posing a long-term threat to ecosystem stability and biodiversity (Alvarado et al., 2020), and it is these ecosystems where anthropogenic climate forcing is most likely causing increase in burning.~~

For regional BA totals, the likelihood that socioeconomic drivers increased BA was 47% (**Figure 14**), indicating no clear signal that human landscape modification influences the extent of burning in seasons like early 2024. The estimated Amplification Factor was 1.08, but with a wide 90% confidence interval of 0.44 to 7.21 (**Table 5**). The wide confidence range, from potential halving of BA to a seven-fold increase, indicates that our model finds socioeconomic drivers to have a highly uncertain influence on regional fire activity during this period. This uncertainty likely reflects both the limited resolution of the socioeconomic variables used (e.g. population density, broad land cover classes) and the challenge of capturing the complex ways that human activities interact with fire. It is also possible that opposing effects such as suppression in one area versus ignition pressure in another, could be offsetting each other in regional statistics, though the modelling framework does not resolve these interactions explicitly.



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Figure 14: Probability density of the Amplification Factor (AF) for each region, showing how different factors influenced the extent of burning for each focal region. The top panel displays results for the entire region, while the bottom panel focuses on sub-regional extremes (defined as the grid cells in the top 5% of BA fraction). Anthropogenic climate forcing targets the 2024/25 focal moths using NRT set up with counterfactuals using all HadGEM ensemble members; socioeconomic factors uses the ISIMIP set up, looking at increased likelihood of 2024/255 like events in 2003-2019 with climate trends removed vs 1900-1917. An AF greater than 1 indicates that the factor contributed to an increase in burned area extent; a AF less than 1 indicates a reducing influence; a value near 1 suggests no change. Vases show probability distribution of AF, dots within each vase show central estimate and bars show 90th percentile confidence range. The percentages lower left of each vase shows the likelihood of each factor increasing burned area.

5.2.2.2. Pantanal and Chiquitano

The Pantanal and Chiquitano regions showed one of the strongest anthropogenic climate change signals of all focal regions studied here or in previous reports (Jones et al., 2024b). The likelihood that anthropogenic climate forcing increased the observed regional BA is estimated at 88% (**Figure 14**), indicating anthropogenic climate forcing likely drove an increase in BA (**Table 5**). The total BA was 34.5 times higher (our *amplification factor*) in the factual ensemble than in the counterfactual, although the wide uncertainty range of 0.84 to 100 suggests the effect of anthropogenic climate change could range from minimal to extremely large (**Table 5**). When internal meteorological variability is removed (using ensemble-mean), the estimated amplification factor remains largely unchanged. The model-based risk ratio for the event is 3.3, meaning the observed extent was roughly three times more likely due to anthropogenic climate change.

Climate influence was relatively consistent across the region (**Figure 15**). Uniformity in attribution results may reflect the broad scale influence of anthropogenic climate change. It also suggests that climate change is amplifying fire risk even in areas with relatively intact ecosystems or seasonal wetlands, underscoring the vulnerability of these landscapes to the ongoing warming. However, the wide range in uncertainty highlights the need for improved observational data and better representation of fuel-moisture dynamics in fire-prone wetland mosaics such as Pantanal.

Historically, the Pantanal has been characterised by limited human fire use and relatively low natural fire incidence, largely due to its wetland-dominated landscape and seasonal flooding that constrained fuel ignitability (e.g. Alho et al., 2019). This makes the emergence of a clear climate-change signal particularly notable: in a system where fire has not been a dominant disturbance, anthropogenic drying is now creating conditions conducive to burning. The broader Pantanal-Chiquitano region, however, also includes drier savanna and forest mosaics where fire use has been more common, meaning that the 2024 anomaly reflects both the novel susceptibility of wetlands and the intensification of burning in adjacent ecosystems. This historical context strengthens the case that, in a region that has not traditionally experienced large-scale burning the 2024 fires arose from compounding drivers: climate change removing long-standing hydrological constraints on fire and human activity providing ignition and land-use pressures.

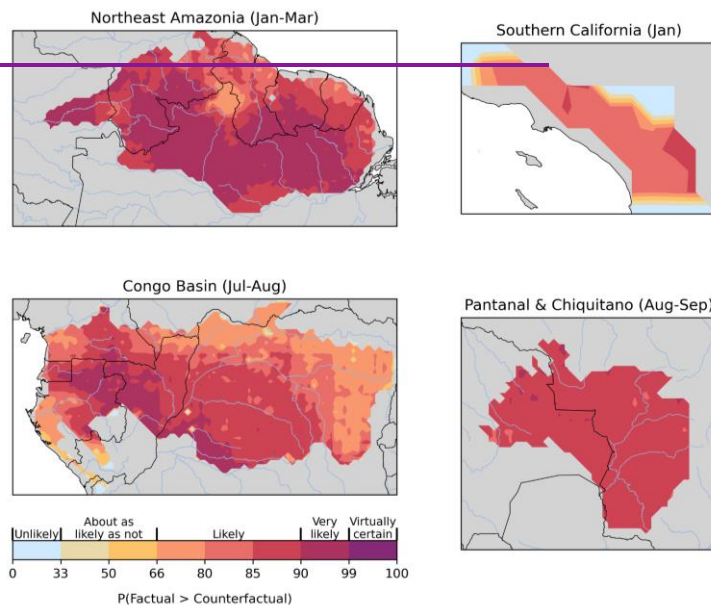
Socioeconomic factors show a very strong role for direct human influence in shaping BA anomalies during 2024-like events in the Pantanal and Chiquitano region. At the regional scale, the likelihood that socioeconomic factors increased BA is 99%, with an estimated amplification factor (AF) exceeding 100 (90% confidence interval: 2.12 to 100). This means that even under conservative estimates, human activity at least doubled BA during comparable fire years. In sub-regional extremes, the Amplification Factor range is even more extreme with a central estimate of more than 100 (lower 90% confidence bound of 16.24), with a similarly high likelihood (>99%) that human activity contributed. This implies that the vast majority of burning in these most severely affected areas was directly linked to socioeconomic drivers and would have been extremely unlikely in their absence.

These results confirm that direct human influences, such as land use effects and human ignition sources, can be as significant, or more so, than climate change in raising the likelihood of extreme wildfire events in the Pantanal-Chiquitano region. This is particularly important and promising because these factors can be directly addressed through local policies, incentives and enforcement actions, offering clear and locally isolated pathways for intervention and risk reduction alongside global action on climate change. Our socio-economic analysis is necessarily based on relatively coarse indicators (population density and land-use change), and so should be viewed as a restricted but internally consistent representation of direct human influence. Nonetheless, the framework captures a wide range of plausible

3118 anthropogenic effects and indicates a substantial role for people in shaping fire outcomes.
3119 Importantly, this finding is consistent with independent evidence from recent fire extremes in
3120 the Pantanal, including the 2020 event (Barbosa et al., 2022; Barbosa, 2024), which shows
3121 that land management, ignition patterns, and water extraction have amplified fire risk
3122 alongside climate pressures. Previous studies highlight that management of ignition sources,
3123 enforcement of land use regulation, and reduction of wetland degradation and water extraction
3124 can reduce fire vulnerability in this region (Barbosa et al. 2022; Menezes et al. 2022 ,Marques
3125 et al 2021). Thus, while global mitigation remains essential, locally targeted actions represent
3126 concrete and tractable pathways for reducing future fire risk. T The consistency between
3127 regional and sub-regional attribution indicates that these influences are not just diffuse but are
3128 concentrated in areas of greatest impact. Even the lower bounds of the confidence intervals
3129 provide compelling evidence that anthropogenic pressure substantially elevated fire
3130 outcomes.

3131
3132 These results agree with a growing body of evidence pointing to compounding non-linear
3133 effects of human and climatic drivers in the Pantanal (Marques et al., 2021, Barbosa et al.,
3134 2022, Santos et al., 2024). While this attribution includes some of the human drivers identified
3135 in the region, such as land use change, other key drivers, like wetland degradation and water
3136 extraction (which can intensify fire risk by drying out the landscape; Barbosa et al., 2022,
3137 2025b), are not captured here.

3138



3139

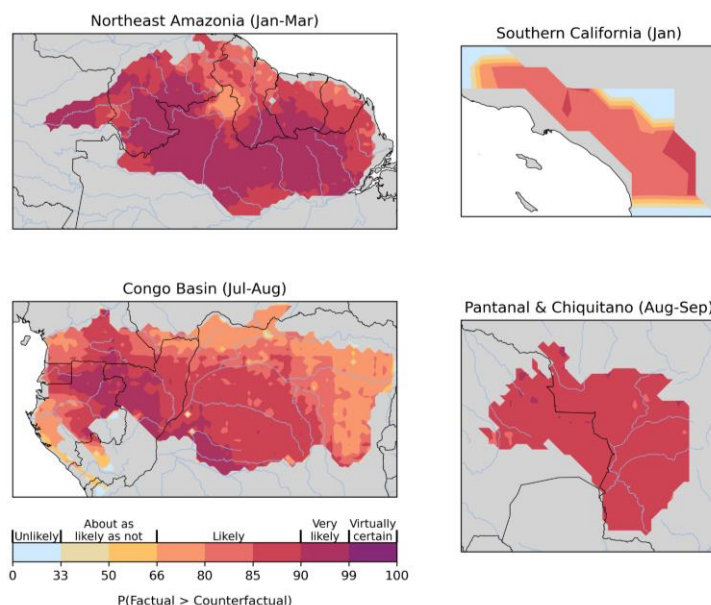


Figure 15: Regions where anthropogenic climate forcing most likely influenced fire activity during the 2024-25 fire season, based on the ConFLAME Near Real-Time setup. Maps show the probability that burned area (BA) was higher in the factual (climate change-influenced) scenario compared to the counterfactual (no climate change) scenario-based on the proportion of ensemble members where BA was greater in the factual than in the counterfactual scenario. Results are shown for focal fire periods in each region: January-March 2024 for Northeast Amazonia; August-September 2024 for the Pantanal and Chiquitano; June-August 2024 for the Congo Basin; and January 2025 for Southern California. Colourbar descriptive labels are based on IPCC uncertainty definitions (Mastrandrea et al. 2010).

5.2.2.3. Southern California

Anthropogenic climate forcing likely contributed to the high levels of BA observed in Southern California in January 2025, with a likelihood of increased burning of 89%. The amplification factor (AF) was estimated at 24.8, though with a wide uncertainty range (90% confidence interval 0.89 to 100), indicating that the influence could have ranged from negligible to extremely large. Despite this spread, the ensemble-mean counterfactual results largely agree, reinforcing confidence that anthropogenic climate forcing increased the likelihood of the event. The risk ratio of 2.3 suggests that similar fire [weather](#) conditions are more than twice as likely in the present-day climate compared to a scenario without climate change. This elevated risk was in January, outside the region's typical peak fire season, suggesting that anthropogenic forcing may be expanding the seasonal window during which large fire events can occur.

There is no clear evidence [from our analysis](#) that socioeconomic factors occurring on the landscape increased the likelihood of January 2025-like regional BA in Southern California during 2002-2019. The estimated likelihood of an increase is 55%, with a highly uncertain amplification factor (AF = 1.04 [0.17-85.58]). As with the climate attribution, this likely reflects the small size of the region and limited signal in long-term data.

5.2.2.4. Congo Basin

Anthropogenic climate forcing likely increased the total area burned across the Congo Basin during June to August 2024. The likelihood of an increase is estimated at 92%, with an amplification factor (AF) of 2.69, meaning the event-scale BA was nearly three times higher than it would have been without forcing. However, there remains some uncertainty: while the best estimate points to a substantial increase, the range spans from a very small influence to a more than 30-fold increase (90% confidence range of 0.96 to 33.96).

When we account for internal climate variability by averaging across all ensemble simulations (rather than using only the observed event conditions), the signal strengthens substantially. In this case, anthropogenic climate change appears to have increased BA by a factor of 15 (90% confidence range: 0.97 to over 100), with a risk ratio of 2.6, which shows a more consistent pattern of increased fire risk due to long-term warming and drying trends. Unlike other regions, where most of the uncertainty stems from how fire responds to environmental conditions, in the Congo Basin uncertainty in the meteorological response to climate change itself plays a larger role.

The influence of climate change also varied significantly within the region. The strongest signal appears in the southern parts of the Congo Basin, particularly the Southern Moist Forests, where our modeling frameworks suggest climate change very likely (90-95% likelihood), using IPCC terms definition **Figure 15** increased BA. Further north, in the DRC's northern moist forests, the likelihood was lower (50-80%), and in the Southern Gabon transition forests, there was little to no signal. These spatial differences may reflect varying sensitivities to rainfall patterns, fuel conditions, or other landscape features, and highlight the importance of region-specific analysis.

There is no clear signal in our indicators (population density and land-use change) that socioeconomic factors increased BA during the June–August 2024 fires in the Congo. ~~Across the region as a whole, the Basin. The~~ likelihood of increased burning ~~due to population density and land-use change~~ was 26%, ~~with an amplification factor~~ regionally (AF) ~~of = 0.94 (90% confidence interval CI: 0.70 to 1.17),~~ suggesting a small or even slightly dampening influence. At the sub-regional level, attribution remains uncertain. ~~The, with estimates of 62% likelihood of increased BA in the most affected grid cells was estimated at 62%, (AF = 1.00 [0.68–1.69]).~~ These results should be interpreted cautiously, as our indicators cannot capture the full dynamics of the region. For instance, population density alone does not represent how local conflicts (Meddour-Sahar, et al. 2013; Trigg et al. 2011) shifting cultivation practices vary with an AF of 1.00 [0.68–1.69].

3209 [fallow length or sedenterisation \(Molinario et al. 2020\); agricultural land-use fractions capture](#)
3210 [broad changes in cover but not day-to-day fire use for field preparation; and land-cover change](#)
3211 [highlights large-scale transitions but not how governance or local management influence](#)
3212 [burning \(Tyukavina et al. 2018 ; Perkins et al., 2024\). In the Congo, these processes interact](#)
3213 [in complex and sometimes opposing ways, shorter fallow periods can increase fire use, while](#)
3214 [sedenterisation can reduce it \(Perkins et al., 2024\). Incorporating such dynamics into future](#)
3215 [frameworks will be important for improving and tightening socio-economic attribution in this](#)
3216 [region.](#)

3217 **Table 5:** Summary of attribution results for burned area (BA) and fire weather indices during key fire events across Northeast Amazonia (Jan-
 3218 Mar 2024), Pantanal-Chiquitano (Aug-Sep 2024), Congo Basin (Jul 2024), and Southern California (Jan 2025). Values are reported for both
 3219 burned area (BA) across the full region and sub-regional extremes -the areas that saw the most burning (see Figure 5). Metrics include the
 3220 amplification factor (AF; the ratio of BA under the influence of the assessed factor relative to the counterfactual), risk ratio (RR) of fire weather
 3221 index during the events, and % change in annual mean (background) BA. Results are shown for different configurations: anthropogenic
 3222 meteorological forcing (using near real-time and ensemble-mean setups), total climate forcing, and socio-economic factors. Values are reported
 3223 as median [5th-95th percentile] ranges, with likelihoods indicating the confidence that the factor contributed to increased burning or extreme fire
 3224 weather. Colours indicate IPCC-defined confidence or likelihood categories (Mastrandrea et al., 2010). Where likelihoods are not explicitly
 3225 provided, colours reflect the lowest plausible category based on the reported confidence range.

3226

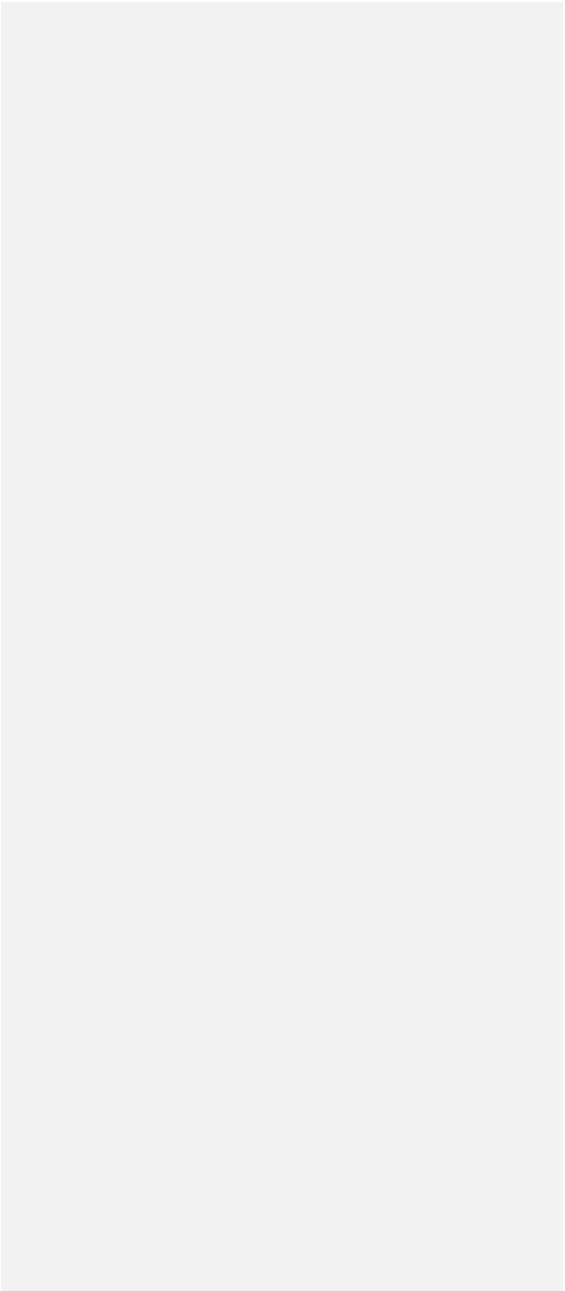
Variable	Metrics	Sources	Northeast Amazonia	Pantanal and Chiquitano	Southern California	Congo Basin
Anthropogenic climate forcing						
Fire Weather Index	Risk Ratio (RR)	HadGEM	31.96-72.64	4.16-5.45		3.04-8.00
	RR	CanESM5	1.9 [1.5, 53.3]	12.3 [3.4, 76.9]	1.7 [1.6, 1.8]	1.3 [0.7, 1.7]
	RR	WWA		4.6 [1.1, 20]*	1.37 [0.48, 3.6]+	
	Intensity Delta	WWA		+39% [13%, 71%]*	+5.7% [-10, 27]+	
Burned Area (BA)	Amplification factor (AF)	ConFLAME/ HadGEM ensemble	4.33 [1.02, 25.32]	34.47 [0.84, >100]	24.79 [0.89, >100]	2.69 [0.96, 33.96]
	RR		2.1	3.3	2.3	1.6
	AF	ConFLAME/ HadGEM mean	13.25 [1.02, >100]	>100 [0.82,- > 100]	>100 [0.82, >100]	14.76 [0.97, > 100]
	RR		2.7	3.5	2.9	2.6
Areas of highest BA	AF	ConFLAME/ HadGEM ensemble	1.17 [1.01, 5.13]	1.91 [0.98, >100]	>100 [0.95, >100]	1.29 [0.96,-3.32]
	R		2.2	2.4	2.9	1.8

	AF	ConFLAME/ HadGEM mean	1.11 [0.95,1.94]	>100 [0.96, >100]	>100 [0.98, >100]	1.24 [0.84,-1.55]
	RR		1.6	3.7	3.8	1.3
Variable	Metrics	Sources	Northeast Amazonia	Pantanal and Chiquitano	Southern California	Congo Basin
Total climate change						
BA	AF	ConFLAME/ ISIMIP	1.01 [0.88,1.15]	>100 [2.73, >100]	1.07 [0.68, 2.83]	1.08 [0.95, 1.43]
Areas of highest BA	AF	ConFLAME/ ISIMIP	1.02 [0.94,1.13]	>100 [4.92, >100]	1.00[0.91, 1.86]	1.14 [0.87, 3.02]
Background BA		FireMIP	-6% [-11%, - 2%]	10% [6, 15%]	7% [2%, 12%]	54% [45%, 63%]
Socio-economic factors						
Burned Area	AF	ConFLAME/ ISIMIP	0.99 [0.8, 1.41]	>100 [2.12, >100]	1.04 [0.17, 85.59]	0.94 [0.7, 1.17]
Max. Burned Area	AF	ConFLAME/ ISIMIP	1.02 [1.07, 1.13]	>100 [16.24, >100]	1.00 [0.85, 6.,65]	1.00[0.68, 1.69]
Background Burned Area		FireMIP	10% [3%, 17%]	-7% [-12%, -2%]	-3% [-7%, -1%]	-16% [-21%, -11%]
All forcings						
Burned Area	AF	ConFLAME/ ISIMIP	0.99 [0.81, 1.47]	1.08 [0.44, 7.21]	1.05 [0.26, 64.26]	1.01 [0.86, 1.42]
Max. Burned Area	AF	ConFLAME/ ISIMIP	1.01 [0.96, 1.10]	1.04 [0.98, 8.26]	1.00 [0.86, 12.16]	1.06 [0.73-4.44]
Background BA		FireMIP	1% [-6%, 9%]	3% [-2%, 9%]	2% [-2%, 7%]	25% [18%, 33%]

Virtually certain	Very likely	Likely	About as likely as not	Unlikely,	Very unlikely	Exceptionally unlikely
> 99%	>90%	>66%	33-66%	<33%	<10%	<1%

|

3228 *WWA results for June DSR and
3229 + January max FWI



5.2.3. Background Changes in Burned Area this Century

We assess how climate and socio-economic drivers have influenced changes in background levels of BA for each focus region using the global fire model attribution framework introduced by Burton, Lampe et al. (2024), adapted this year to match the specific geographic areas analysed in this report (see methods in **Section 5.1.4**). Results represent the change in median monthly BA during 2003-2019 compared to a counterfactual scenario in which anthropogenic climate change or changes in socio-economic factors were removed. This is distinct from our analyses focussing on the attribution of individual focal events in **Sections 5.2.1 and 5.2.2**.

5.2.3.1. Northeast Amazonia

Total climate forcing led to a modest but consistent decrease in background BA between 2003-2019, with a median change of -6% [-11%, -2%] compared to a counterfactual without climate change. Unlike the earlier attribution method (**Section 5.2.2**), which focused on extreme 2024-like events, this model captures long-term, background fire activity, including broader fuel-climate interactions.

The reduction in BA may reflect increased moisture or changes in vegetation structure that reduce flammability, though the exact mechanism is unclear. Recent observational analyses suggest a rise in wet-season (December to May) rainfall and a reduction in dry days in northern Amazonia over the past two decades (Barichivich et al., 2018; Almeida et al., 2017), which could contribute to these trends if captured in the climate inputs. The underlying models used in this attribution framework used here also features tighter coupling between vegetation, climate, and fire than the event-based approach, which may explain some of the differences, though it remains difficult to determine whether these are due to improved fuel representation or simply reflect a contrast between background and extreme conditions.

Socioeconomic changes are estimated to have increased the background BA in Northeast Amazonia by +10% [3%, 17%] in 2003-2019 compared to 1901-1917. This signal aligns well with the earlier analysis of 2024-like events (**Section 5.2.2.1**) but is more narrowly constrained, reinforcing the role of human-driven changes as a key influence on regional fire activity, as identified in many previous studies. For instance, recent studies on land use and fire dynamics in the Amazonia region points to rising fire activity associated with expanding agricultural areas, secondary vegetation, and newly deforested areas (Silveira et al., 2022). Human activities remain the primary source of ignition, mainly through practices such as deforestation, pasture maintenance, and crop field burning, often intensified under dry conditions (Lapola et al., 2023).

5.2.3.2. Pantanal and Chiquitano

We find a modest but robust signal of climate-driven change in background fire activity. Between 2003 and 2019, total climate forcing is estimated to have increased the average BA by 10% [6%, 15%]. The relatively narrow confidence range suggests strong model agreement and indicates that the region's area burned has already been measurably affected by long-term climatic shifts. This aligns with broader lines of evidence that highlight the Pantanal's vulnerability to changes in rainfall patterns and dry season intensity, which influence both fuel availability and flammability (**Section 4.2**). These findings are also consistent with attribution results for extreme events in 2024 (**Section 5.2.2.2**), which also showed a high likelihood of increased burning, albeit with greater uncertainty.

We estimate that socioeconomic drivers contributed a reduction in background BA of 7% [-12%, -2%] compared to pre-industrial conditions. This suggests that long-term changes in land

use and management, including shifts in agricultural practices, may have contributed to a modest but consistent suppression of average fire activity over the past two decades. The attribution of socioeconomic influence on BA in the Pantanal presents an interesting contrast with the attribution of focal event BA in the previous section, which suggests that socioeconomic factors very likely increased BA (Section 5.2.2.2). This contrast may point to important temporal and functional differences:

- Long-term socioeconomic changes, such as improved fire control in settled areas or changes in land use, could suppress background fire activity.
- Yet, during extreme conditions, these same systems may fail to contain fires, or different areas (e.g. the interface between private properties and protected areas, Barbosa et al., 2022) may dominate the fire signal.

~~Still, the~~ We estimate that socioeconomic drivers, here represented by mainly land-use, contributed a modest reduction in background BA of 7% [-12%, -2%] compared to pre-industrial conditions. This suggests that long-term changes in land use may have contributed to a slight suppression of average fire activity over the past two decades. It is important to stress that our analysis is limited to land-cover change and population density proxies, and does not capture the full spectrum of socio-economic drivers such as ignition practices, fire suppression, or governance, which may also influence fire activity. The contrast with focal-event BA attribution, which indicates that socioeconomic factors very likely increased BA (Section 5.2.2.2), may therefore reflect differences in timescale, the specific processes captured by our indicators, or areas where our representation of human influence is incomplete. Region wide BA attribution above, for example, also incorporates land-use change and locally optimised, more detailed population density-based representations. However, as noted in Section 5.2.2.2, even there we miss crucial aspects of local human agency, which can vary substantially in space and time across the region (Perkins et al. 2024).

•

This disagreement raises a cautionary flag. While the two methods target different timescales and use different models, their confidence intervals do not fully overlap, suggesting that at least one framework may be underestimating uncertainty or missing key processes. the need for caution in interpreting long-term trends: while our framework provides internally consistent estimates within its scope, it cannot fully resolve the broader socio-economic influences on fire. It also reinforces the importance of using multiple, independent lines of evidence in attribution work and, specifically for the Pantanal, shows that more work is needed to assess the balance between human impact on background vs extreme BA along with the modelling techniques used to assess this.

5.2.3.3. Southern California

In Southern California, the models attribute a +7% [2%, 12%] increase in median background BA to total climate forcing. This is consistent with the attribution results for 2025-like events (Section 5.2.2.3), though with higher confidence. The agreement across these distinct approaches, despite targeting different fire outcomes (seasonal extremes vs general background activity), provides additional confidence that long-term climate change is influencing baseline fire conditions in the region.

Socio-economic influences contributed a -3% change in background BA, with an uncertainty range of [-7%, 1%]. While not statistically significant, this result is more tightly constrained than those from the earlier analysis of 2025-like events. The modest downward influence may reflect intensifying suppression capacity, declines in human-caused fires due to fire-prevention

3336 policies including those targeted to electrical utilities (Jorge et al., 2025; Abatzoglou et al.,
3337 2020), or other urban interface factors, though uncertainty remains high.

3338 3339 **5.2.3.4. Congo Basin**

3340
3341 In the Congo Basin, we estimate that total climate change has driven an increase in mean
3342 annual BA of 54%, with a tight confidence range of [45%, 63%]. This makes it one of the most
3343 robust signals of climate influence across the background fire analyses. These results are
3344 consistent with, though slightly stronger and more confident than, the attribution using 2024-
3345 like extreme events. The agreement between methods strengthens confidence that climate
3346 change is already amplifying baseline fire activity in the region.

3347
3348 This signal likely reflects a clear climate influence on fire-conducive weather, particularly in
3349 the southern part of the basin (**Section 4.2.2.2.4**). While fuel limitations played a role in
3350 moderating fire spread (**Figure 12**), the background increase in BA appears strongly tied to
3351 meteorological shifts linked to climate change.

3352 Socioeconomic influences appear to have played a moderating role in background fire activity
3353 across the Congo Basin. In our process-based model analysis, socioeconomic drivers,
3354 including changes in land use, land cover, and population, led to a 16% reduction in
3355 background BA between 2003-2019, with a 90% confidence range of -21% to -11%. This
3356 suggests a consistent and substantial dampening effect on fire, possibly reflecting a
3357 combination of land fragmentation, land use conversion, or reduced fire use. These results
3358 are broadly in line with, though more confidently constrained than, the amplification factor
3359 estimated for 2024-like events in the previous attribution method, which indicated limited
3360 influence from socioeconomic factors.

3361 **6. Seasonal and Multi-Decadal Outlook**

3362 3363 **6.1. Methods**

3364 3365 **6.1.1. Seasonal Forecasts**

3366 3367 **6.1.1.1. Fire Weather Index**

3368 In **Section 4**, we introduced the use of seasonal forecasts of FWI and examined how they
3369 performed during the focal events of the 2024-25 fire season. In this section, we present global
3370 FWI forecasts from the ECMWF's SEAS5 seasonal prediction system for the months June-
3371 August 2025, extending the same approach employed in **Section 4** throughout the boreal
3372 summer months of 2025 (see **Section 4.1.1.2.1** for methods).

3373 **6.1.1.2. Burned Area**

3374 In **Section 4**, we introduced the use of seasonal forecasts of burned areas using a
3375 combination of weather driver and ML and examined how they performed during the focal
3376 events of the 2024-25 fire season. In this section, we present global BA forecasts from the
3377 same system for the months July-September 2025, extending the same approach employed
3378 in **Section 4** throughout the boreal summer months of 2025 (see **Section 4.1.1.2.2** for
3379 methods).

3380 6.1.2. Multi-Decadal Projections

3382 6.1.2.1. Fire Weather Index at Future Global Warming Levels

3384 To calculate how the risk of fire weather extremes might evolve with future warming, we apply
3385 the same framework described in **Supplementary Material S5.1.1** but instead of comparing
3386 recent climate to the past, we compare it to a set of global warming levels: 1.5 °C, 2.0 °C, 3.0
3387 °C, and 4.0 °C above recent past climate (2016-2025).

3388
3389 For each level of warming, we identify years in the CanESM5 ensemble where the smoothed
3390 11-year running global mean temperature aligns with the target level, and then assess the
3391 frequency of extreme 7-day FWI events in those years, as per Liu et al. (2023b) and similar to
3392 Otto et al. (2018). Comparing this to the 2016-2025 climate baseline gives us a forward-looking
3393 set of Risk Ratios (RR) — RR1.5, RR2.0, etc. These indicate how much more likely such
3394 extremes become as the planet warms.

3395
3396 As with the attribution to past climate (**Section S5.1.1**), uncertainties are captured through
3397 bootstrapped confidence intervals, enabling meaningful comparison of future risks even when
3398 rare extremes are involved.

3399 6.1.2.2. Burned Area in Future Emissions Scenarios

3400
3401 In order to project future changes in BA, we extended the ConFLAME ISIMIP3a modelling
3402 approach used in **Section 5.1.3** to future decades under Shared Socioeconomic Pathway
3403 (SSP) scenarios SSP126, SSP370, and SSP585, following a similar protocol to UNEP
3404 (2022a). We use the same optimised model as in **Section 5.1.3**, but here we employ bias-
3405 corrected global climate model (GCM) outputs from ISIMIP3b (Frieler et al. 2025) for
3406 prediction. While ISIMIP3a uses reanalysis data for historical analysis, ISIMIP3b employs
3407 GCM data to project future climates and is designed for usage cases requiring a seamless
3408 continuation of the historical period into future scenarios.

3409
3410 ISIMIP3b utilizes five bias-corrected GCMs, including historical model output up to 2014 and
3411 future scenarios from 2015-2100 under the three SSPs. ISIMIP3b uses surface-based
3412 meteorological outputs from ScenarioMIP simulations, which include future forcings from
3413 greenhouse gases, aerosols, land-use change, and short-lived climate forcers. The five GCMs
3414 used are: GFDL-ESM4 (Held et al., 2019), IPSL-CM6A-LR (Boucher et al., 2020), MPI-ESM1-
3415 2-HR (Mauritsen et al., 2019), MRI-ESM2-0 (Yukimoto et al., 2019), and UKESM1-0-LL (Tang
3416 et al., 2019; Sellar et al., 2019). As part of ISIMIP3b, each GCM is bias-corrected as described
3417 in Lange (2019).

3418
3419 Future ISIMIP3b projections for socioeconomic drivers such as population density or land use
3420 change were not available at time of analysis. As such, our simulations exclude future changes
3421 in ignition sources or direct land-use modification on both fire and vegetation. To simulate
3422 vegetation structure and fuel availability, the JULES-ES dynamic vegetation model was run
3423 offline, driven by surface climate variables from each of the five bias-corrected GCMs under
3424 each SSP scenario, and scenario-specific CO₂ concentrations to represent CO₂ fertilization,
3425 along with prescribed nitrogen deposition but excluding changes in fertiliser application, along
3426 with prescribed nitrogen deposition but excluding changes in fertiliser application. The land
3427 cover output from JULES-ES was then bias-corrected (using the same mapping procedure as
3428 **Section 5.1.3**, based on biases between JULES-ES driven by reanalysis and VCF
3429 observations) to maintain consistency with the GCM bias-correction procedures. Our
3430 approach provides a probability distribution of future BA representing the uncertainty range
3431 from cross-model (GCM) spread in the response of climate and vegetation to emissions for
3432 each scenario and year in the period 2010-2100. Years 2010-2014 were adopted from the
historical experiment for each GCM, and post-2014 from branched SSP and model specific

3433 projections. We describe future changes as significant if the range across GCM projections
3434 for a future period does not overlap with the range given by the GCMs for 2010s.

3435 Using this driving data, we generate 1,000-member ensembles for each region and each
3436 GCM/SSP combination, using the trained ConFLAME-ISIMIP model described in **Section**
3437 **5.1.3**. For each 10-year period, we calculate the likelihood of extreme fires by determining the
3438 fraction of years within each ensemble member where burned area during the event months
3439 exceeds that of the observed focal event. We then average this exceedance fraction across
3440 all 1,000 ensemble members to estimate the likelihood for that decade. This process is
3441 repeated for each GCM and SSP.

3442 For decades beyond 2010s, we then calculate the increase in the likelihood of 2024/25-like
3443 events by taking the ratio of the exceedance frequency in each future decade relative to the
3444 2010s baseline. This is analogous to the risk ratio used in **Section 4**, where the future period
3445 acts as the “factual” and 2010s as the “counterfactual” baseline. Following methods outlined
3446 in **Section 4**, we perform this analysis for the entire region and for “sub-regional extremes” -
3447 the grid cells with the top 5% of BA.

3448 Lastly, we calculated the integrated probability of experiencing a fire event of similar
3449 magnitude to our target region within the expected lifespan of a citizen born in 2023 (the year
3450 of the latest estimate). According to UN population statistics (United Nations Population
3451 Division, 2023), life expectancy at birth is 75.8 years for Brazil, 79.3 years for the USA, and
3452 61.9 years for the Democratic Republic of the Congo (DRC). While the Northeast Amazonia
3453 and Congo Basin regions span multiple countries, most fire anomalies in these regions
3454 occurred in Brazil and the DRC, respectively (**Figure 5**). To account for years beyond 2100 in
3455 the life expectancy of Brazil and the USA, we extrapolated the annual trend in event
3456 probabilities. The integrated probability is calculated as one minus the product of the annual
3457 probabilities of not experiencing a fire event like the focal event, across each year from 2025.

3458 **6.2. Results**

3459 **6.2.1. Seasonal Forecasts of Fire Weather Index and Burned Area Anomalies**

3461 As of mid-2025, neither La Niña nor El Niño conditions are present in the tropical Pacific.
3462 Instead, the climate system has entered an ENSO-neutral phase, according to the latest report
3463 from the National Oceanic and Atmospheric Administration (NOAA, 2025c). This neutral
3464 phase is expected to persist through the remainder of summer, and into at least early autumn.
3465 While neutral ENSO conditions typically indicate a reduced influence of Pacific sea surface
3466 temperature anomalies on global weather patterns, the persistence of anomalously warm
3467 ocean conditions and other climate drivers may continue to exert significant influence on
3468 regional and global climate variability in the months ahead (Frölicher and Laufkötter, 2018).

3469 May 2025 was the second-warmest May on record globally, with an average temperature of
3470 15.79 °C, 0.53 °C above the 1991-2020 climate and 1.4 °C above pre-industrial levels
3471 (Copernicus Climate Changes Service, 2025). While this marked a brief drop below recent
3472 consecutive months exceeding 1.5 °C from pre-industrial record, it still reflects the persistent
3473 trend of global climate warming (Horton, 2025). Unusually low rainfall and soil moisture across
3474 northwestern Europe, including the UK, reached their lowest levels since 1871. This raises
3475 serious concerns about crop failures, potential water shortages and wildfire risk (European
3476 Commission Joint Research Centre, 2025; UK Environment Agency, 2025). Similar conditions
3477 were reported in the US, particularly across Arizona and Texas, where exceptional drought
3478 levels led to reservoir depletion, strict water restrictions, and increased wildfire activity
3479 (National Centers for Environmental Information, 2025; National Interagency Fire Center,
3480 2025).

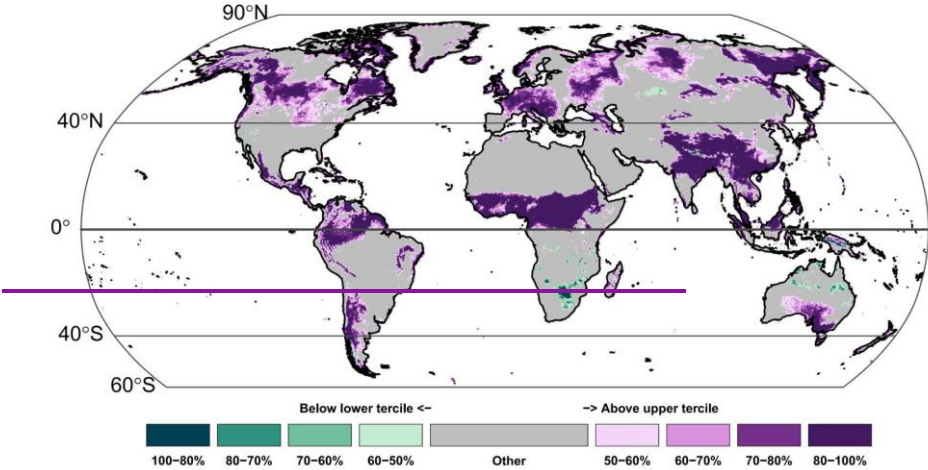
3481 Starting from May, and according to the outlook for the Northern Hemisphere boreal summer
3482 of 2025 (June-July-August), anomalous fire weather conditions are anticipated across several
3483 key regions with high levels of confidence (in places reaching 80 %). Anomalous fire danger
3484 season is expected in Canada, US western states (also see National Interagency Fire Center,
3485 2025), northeast Europe (notably the UK), and parts of Siberia (**Figure 16**). In the equatorial
3486 zone, persistent dryness and hydroclimatic anomalies are expected to increase fire danger
3487 (confidence level of 60% and higher) in Northeast Amazonia, the Congo Basin, and the
3488 Himalayan foothills (affecting areas of India and Nepal). In contrast, a relatively quiet fire
3489 season is projected for the Southern Hemisphere, with only Chile and southern Australia
3490 showing fire-prone conditions at a moderate level of confidence (>50%).

3491 The BA anomaly forecast (bottom panel of **Figure 16**) displays a distinct pattern from that of
3492 FWI, as it models the expected fire response conditioned on both coincident and antecedent
3493 climate variables, based on region-specific statistical relationships. For instance, elevated
3494 probabilities of above-median BA are projected in the western part of South America, southern
3495 California, localized areas of Central America, and central North America. In central Asia,
3496 medium-to-high probabilities emerge, particularly in the eastern regions. In Africa, significant
3497 signals are observed over the central continent, while in Australia, elevated probabilities are
3498 mainly found in the northern regions. Over central Europe, despite a high FWI forecast, limited
3499 historical fire activity prevents reliable calibration of the climate-fire model, and therefore no
3500 BA forecast is issued for this region.

3501

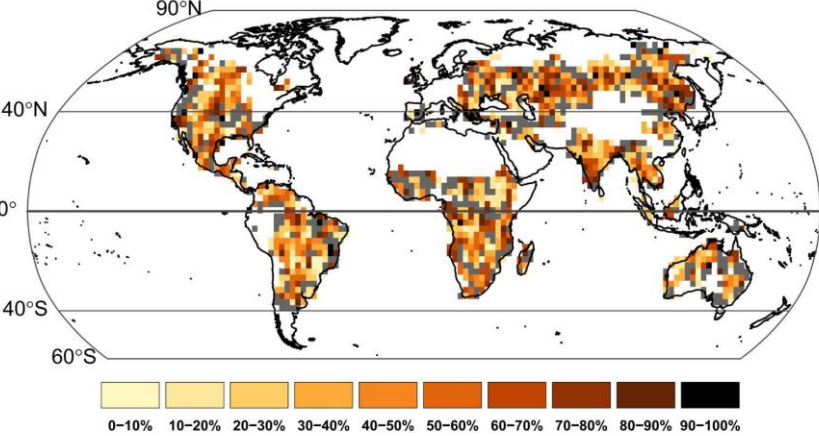
FWI – Seasonal Forecast

Probabilities Below/Above Normal (Start date: 01/06/2025 | Season June–July–August [JJA])



Burned area anomaly – Seasonal Forecast

Probability of burned area anomaly (Start date: 01/06/2025 | Season June–July–August [JJA])



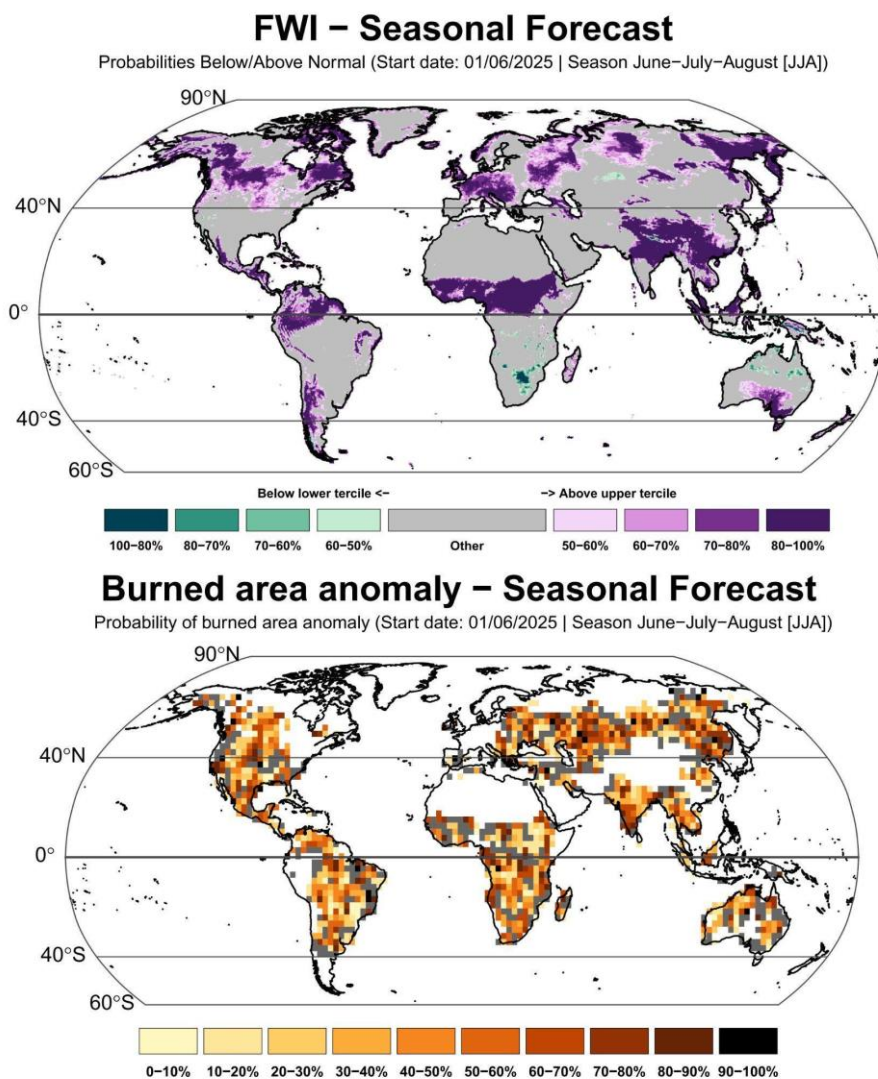
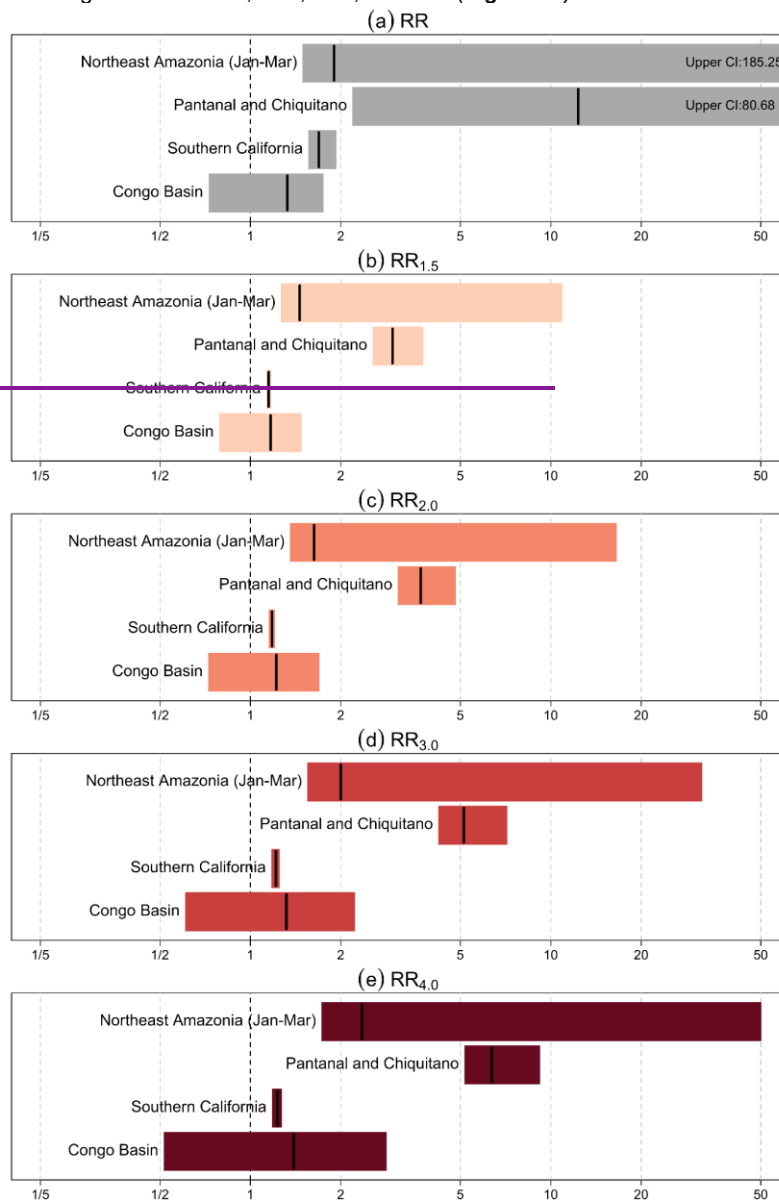


Figure 16: Seasonal prediction of [Fire Weather Index \(FWI\)](#) and [burned area \(BA\) anomalies](#) for the boreal summer of 2025 (June–July–August). Both forecasts are issued in June 2025 and are presented in probabilistic terms: FWI prediction shows the likelihood for increased (above the upper tercile) or decreased (below the lower tercile) fire-weather conditions; whereas BA prediction shows the probability of BA anomalies being above the climatological median. Grey areas are masked where insufficient BA statistics are available to perform the predicted mean.

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6.2.2. Future Changes in Likelihood of Extreme Fire Weather Events

In three of the focal regions where climate change significantly increased the likelihood of a 2024-25-level fire weather event (**Section 5.2.1**), even greater increases are projected under future warming levels of 1.5 °C, 2 °C, 3 °C, and 4 °C (**Figure 17**).



3517

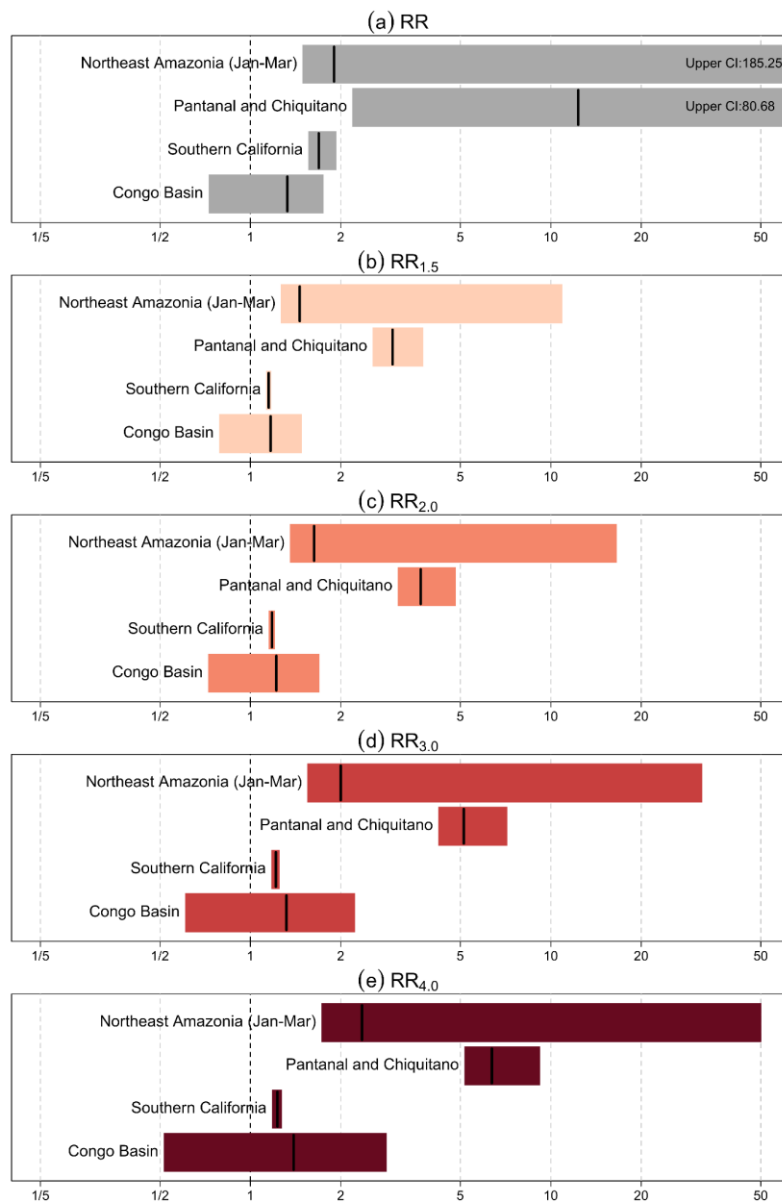


Figure 17: Risk Ratio (RR) estimates based on the comparison between **(a)** the past climate of 1850-1859 and the recent climate of 2016-2025, **(b)** the recent climate of 2016-2025 and the period that global mean surface temperature (GMST) reached **(b)** 1.5 °C, **(c)** 2 °C, **(d)** 3 °C and **(e)** 4 °C for the four extreme wildfire events between 2024 and early 2025 using CanESM5. Bars show 95% confidence intervals (CIs) and central values are shown in bold.

3525 **6.2.2.1. Northeast Amazonia**

3526
3527 In Northeast Amazonia the increased fire weather risk found in **Section 5.2.1.1** during
3528 January-March is projected to continue rising under future warming, with increases in
3529 probability of 1.5 (95% CI: 1.3-10.8), 1.6 (1.4-16.3), 2.0 (1.6-31.4) and 2.4 (1.7-49.5) at 1.5 °C,
3530 2 °C, 3 °C, and 4 °C of warming, respectively . Compared to southern Amazonia, fires in
3531 Northeast Amazonia have gathered less attention from the scientific community and little is
3532 known about how future changes in fire weather conditions may impact this region.

3533
3534 Amazonia spans multiple countries, making coordinated fire governance particularly
3535 challenging. These countries often have differing political priorities and economic interests,
3536 which shape land use policies, enforcement capacity, and investment in fire monitoring and
3537 response systems. Such disparities can hinder the implementation of integrated fire
3538 management strategies, especially in border regions where transboundary fires may occur but
3539 fall under fragmented jurisdictional and institutional frameworks. These institutional and policy
3540 asymmetries introduce further uncertainty about how fire risk will evolve in a warming climate.
3541 As fire weather intensifies, the region's unique fire season and cross-border governance
3542 dynamics should be explicitly considered in fire risk assessments and regional adaptation
3543 strategies.

3544
3545 **6.2.2.2. Pantanal and Chiquitano**

3546
3547 The Pantanal and Chiquitano region, which showed the largest historical increase with 4.75
3548 (95% CI: 4.2-5.5, **Section 5.2.1.2**), is set to continue to increase with global warming, with
3549 projected increases in probability of 3.0 (95% CI: 2.6-3.6), 3.7 (3.2-4.6), 5.1 (4.4-6.5), and 6.4
3550 (5.4-8.3) at 1.5 °C, 2 °C, 3 °C, and 4 °C of warming, respectively (**Figure 17, b-e**). This is
3551 especially concerning for the Pantanal and Chiquitano, where fires are strongly driven by
3552 climate, particularly through extreme (Silva et al., 2022; Barbosa et al., 2022) and compound
3553 events (Ribeiro et al., 2022; Libonati et al., 2022). The ongoing reduction of wetlands in the
3554 Pantanal, often replaced by flammable grasslands (Damasceno-Junior et al., 2021), combined
3555 with the projected increase of fire weather conditions (Feron et al., 2024), may indicate a
3556 permanent shift in the landscape and its fire regime. This increases the vulnerability of fire-
3557 sensitive vegetation and wildlife habitats, while also threatening economic activities that rely
3558 on seasonal flooding.

3559
3560 **6.2.2.3. Southern California**

3561
3562 Southern California shows a similar pattern, with the likelihood of 2024-25 extreme fire
3563 weather being about 1.7 times higher (95% CI: 1.6-1.8) than in the past, and projected
3564 increases in likelihood ranging from 1.1 to 1.3 with rising global temperatures.

3565
3566 **6.2.2.4. Congo Basin**

3567
3568 In contrast, the Congo Basin shows a more modest and statistically non-significant change,
3569 with the likelihood of a similar extreme fire weather event to that of the 2024-25 season
3570 increasing by a factor of 1.3 from the past to the present. Future projections suggest a wide
3571 but uncertain range of change, between 0.5 and 2.7 depending on the warming level.

3572
3573 **6.2.3. Future Changes in Likelihood of Extreme Fire Events**

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6.2.3.1. Northeast Amazonia

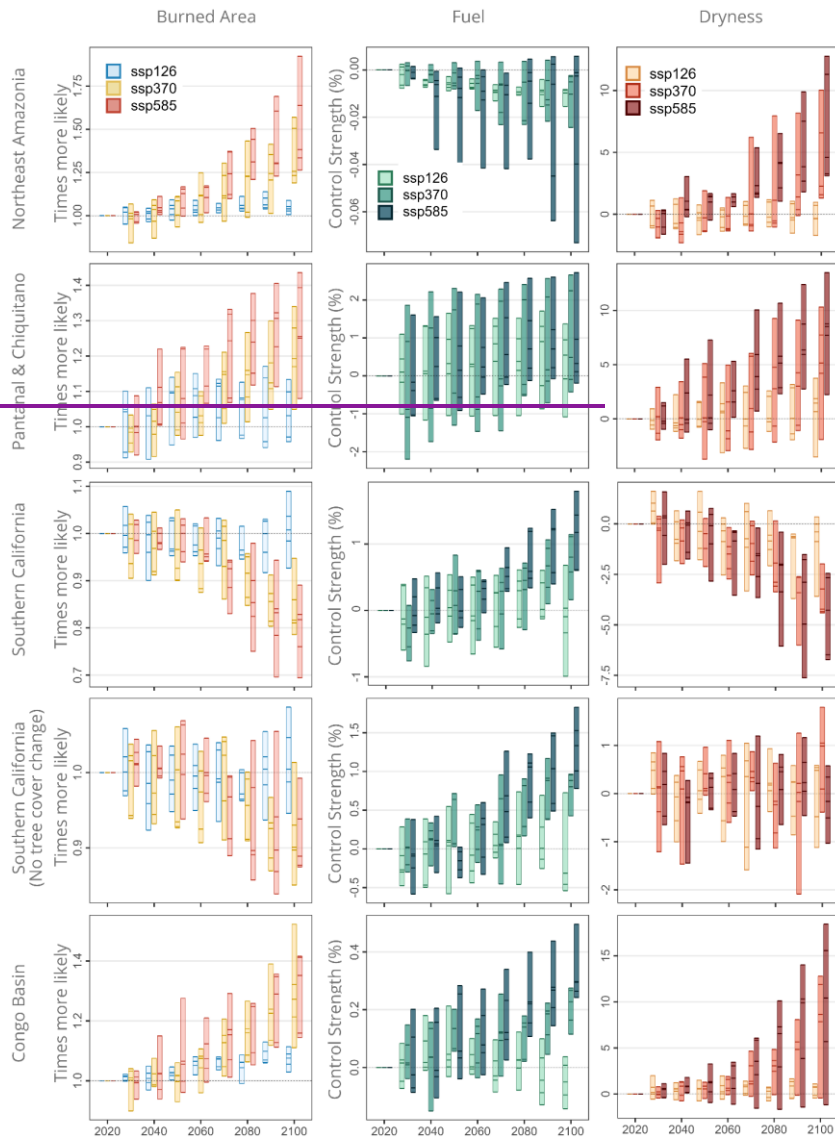
By the 2040s, under SSP585, the likelihood of an event similar to those of the 2024-25 season increases modestly but significantly to 0.12-0.14%, a ~17% increase in frequency compared to the 2010s (**Figure 18; Table 6**). Other scenarios show smaller or even negligible changes over this period. By the end of the century, however, all scenarios project notable increases in event frequency. SSP585 shows the largest rise, with the probability of such an event nearly doubling (up to 1.92 times more frequent). SSP370, reflecting current emissions trajectories, projects a 1.19-1.57 times increase. In contrast, SSP126 illustrates the mitigation potential of low-emission pathways, limiting increases to just 1.09 times (under 10% increase) by 2100, significantly lower than under higher-emission scenarios. SSP370 only clearly diverges from SSP126 by late century (2090s), though the potential for larger increases appears earlier (**Figure 18**). This divergence between the two scenarios is especially pronounced when focusing on areas with the highest BA (top 5% of grid cells, **Figure S29**). These regions of extreme burning could see a doubling in fire extent by mid-century and at least doubling (potentially tripling) by 2100 under SSP370, with substantial overlap with SSP585 projections (where extreme BA could almost quadruple).

By 2100, SSP126 still shows marginal increases in the likelihood of BA events such as those in 2024 (**Figure 18**), though sub-regional extreme BA see much less significant change (**Table 6; Figure S29**), with frequency ranging from slight decreases (by a factor of 0.91) to modest increases (1.34).

These increases are mainly driven by projected declines in moisture availability (**Figure 18 Figure S29**). Although fuel availability is expected to decline somewhat, this only marginally offsets the rise in extreme BA likelihood across the region and has virtually no mitigating effect on the highest BA areas. No changes in fuel are statistically significant in our projections.

Most regions of Northeast Amazonia see increases in January-March (JFM) average BA by 2100 (**Figure S32**). However, under SSP126, increases in the north, French Guiana, Suriname, and Guyana, are less certain and, if they occur, are smaller. This is reflected in a decreased frequency of extremes across these areas (**Figure S32**). Under SSP370, climate change drives widespread increases in BA, with corresponding rises in extremes nearly everywhere except Roraima (Brazil). Most of Brazil and Venezuela are very likely to see increases in BA even under SSP126, with some moist regions showing rises in extremes under SSP126 and widespread increases under SSP370. Results for SSP585 are similar to those of SSP370, with widespread increases in BA and extremes throughout the region. Importantly, increases in extremes begin in some areas in the near future (**Figures S30-31**). By the 2030-2040s, Amapá (Brazil), northern Pará (Brazil), and southern Suriname are projected to experience more frequent extreme BA events and increased BA under the SSP585 scenario (**Figure S30**). Increases in BA are less certain but still likely under SSP370, with mitigation under SSP126 helping to limit these trends.

Finally, we explored what this means for people's lived experience (**Figure 19**). A person born 75.8 years ago (Brazil's current life expectancy) would have had a 33-36% likelihood of witnessing a fire event like January-March 2024 during their lifetime. This suggests that, although anthropogenic changes have increased the likelihood of such fires (see **Section 5**), these events remain far from certain. Even the modest increases in frequency projected under SSP126 would raise that lifetime likelihood to 41-55% for someone born today (i.e., 2025-2021). Under SSP370 (our current path), the chance rises substantially to 52-69%, and under SSP585, to 55-76%. There is also a substantial rise in the probability of experiencing multiple such events within a lifetime, for example, under SSP370, there is a 17-32% chance of seeing two such events, compared to just 6-8% for those born in the 1940s.



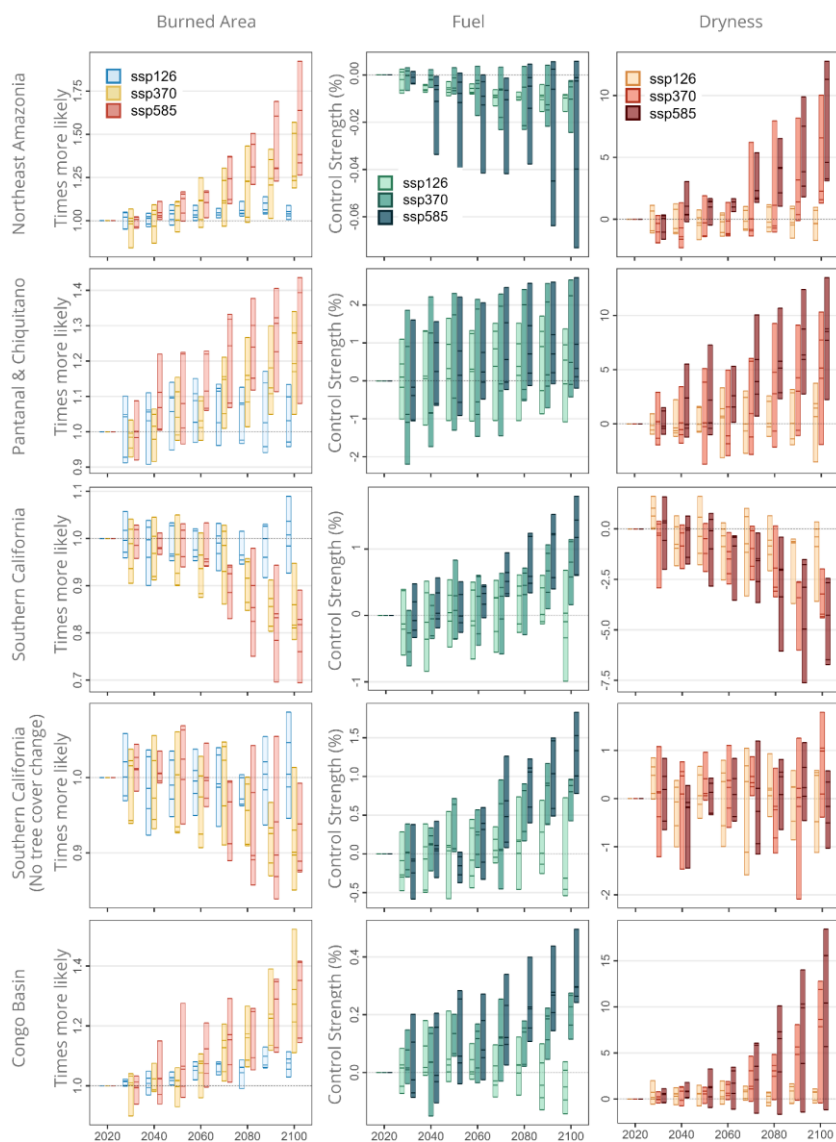


Figure 18: Future projections from ConFLAME of the change in likelihood of BA extent of the magnitude seen in the 2024-25 season, along with the contribution of fuel and moisture conditions in years in which BA exceeds the 2024-25 thresholds. Each set of bars shows changes for each decade relative to the 2010-2020 baseline, with each bar representing a different SSP scenario and the spread of bars indicating the variation across GCMs, with individual bars representing different GCMs.

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Table 6: Summary of the likelihood of extreme events today using reanalysis ‘factual’ and into the future using bias-corrected GCMs for our focal events identified in **Section 4.4.3**. Min and max report range across GCMs. We also determine how much more frequent the events will be at two different time horizons based on each model's likelihood in the future projections over likelihood during 2010-2020. Asterisks (*) indicate non-significant changes from 2010-2020 values. Colours show linear increase of likelihood (orange) and frequency (blue for less frequent, orange for more), where darker shade indicates higher values. The top half of the table displays projections using BA over the entire region, while the bottom shows projections for sub-regional extremes (grid cells with the top 5% BAs).

All Region											
Region	SSP	Represents	Likelihood(%/year during focal months)						How much more frequent (multiplier)		
			2010-2020		2040-2050		2090-2100		2040-2050		2090-2100
			min	max	min	max	min	max	min	max	
Northeast Amazonia	Factual	observed	0.073								
	SSP126	strong mitigation	0.12	0.12	0.12*	0.14*	0.12	0.13	0.97*	1.09*	1.01
	SSP370	~~middle-of-the-road~~ medium-high	0.12	0.13	0.12*	0.13*	0.15	0.19	0.93*	1.11*	1.19
	SSP585	no mitigation	0.12	0.12	0.12	0.14	0.15	0.23	1	1.17	1.26
Pantanal & Chiquitano	Factual	observed	0.19								
	SSP126	strong mitigation	0.09	0.09	0.09*	0.11*	0.09*	0.11*	0.95*	1.14*	0.96*
	SSP370	~~middle-of-the-road~~ medium-high	0.08	0.1	0.09*	0.11*	0.1	0.12	0.98*	1.15*	1.05
	SSP585	no mitigation	0.08	0.1	0.09*	0.11*	0.1	0.13	0.97*	1.22*	1.08
Southern California	Factual	observed	0.38								
	SSP126	strong mitigation	0.34	0.36	0.33*	0.36*	0.31*	0.38*	0.95*	1.03*	0.93*
	SSP370	~~middle-of-the-road~~ medium-high	0.34	0.37	0.3*	0.37*	0.27	0.33	0.9*	1.05*	0.79
	SSP585	no mitigation	0.34	0.35	0.32*	0.36*	0.23	0.31	0.94*	1.03*	0.69
Southern California - no tree cover change	Factual	observed	0.42								
	SSP126	strong mitigation	0.38	0.41	0.38*	0.4*	0.37*	0.42*	0.95*	1.04*	0.95*
	SSP370	~~middle-of-the-road~~ medium-high	0.38	0.42	0.36*	0.41*	0.35*	0.39*	0.93*	1.06*	0.85*
	SSP585	no mitigation	0.38	0.41	0.38*	0.41*	0.34*	0.38*	0.94*	1.07*	0.87*
Congo Basin	Factual	observed	0.17								
	SSP126	strong mitigation	0.16	0.18	0.17*	0.18*	0.17	0.19	1*	1.05*	1.03

	SSP370	middle-of-the-roadmedium-high	0.17	0.19	0.17*	0.19*	0.21	0.26	0.93*	1.06*	1.11	1.52
	SSP585	no mitigation	0.17	0.18	0.17*	0.21*	0.2	0.26	0.96*	1.28*	1.15	1.42
Sub-regional extremes												
Region	SSP	Represents	Likelihood(%/year)						How much more frequent (multiplier)			
			2010-2020		2040-2050		2090-2100		2040-2050		2090-2100	
			min	max	min	max	min	max	min	max	min	max
Northeast Amazonia	Factual	observed	<0.01									
	SSP126	strong mitigation	0.01	0.02	0.01*	0.02*	0.01*	0.02*	0.94*	1.38*	0.91*	1.34*
	SSP370	middle-of-the-roadmedium-high	0.01	0.02	0.01*	0.02*	0.03	0.04	0.92*	1.58*	1.98	3.23
	SSP585	no mitigation	0.01	0.02	0.02	0.02	0.03	0.05	1.15	1.64	2	3.6
Pantanal & Chiquitano	Factual	observed	0.01									
	SSP126	strong mitigation	0.02	0.03	0.03	0.03	0.03	0.03	1.05	1.21	1.03	1.24
	SSP370	middle-of-the-roadmedium-high	0.02	0.03	0.03*	0.03*	0.03	0.04	0.94*	1.23*	1.21	1.45
	SS585	no mitigation	0.02	0.03	0.03*	0.03*	0.03	0.04	0.96*	1.45*	1.26	1.75
Southern California	Factual	observed	0.27									
	SSP126	strong mitigation	0.24	0.26	0.24*	0.25*	0.23*	0.27*	0.96*	1.03*	0.94*	1.12*
	SSP370	middle-of-the-roadmedium-high	0.24	0.26	0.22*	0.27*	0.2	0.24	0.91*	1.08*	0.82	0.97
	SSP585	no mitigation	0.24	0.26	0.23*	0.25*	0.18	0.23	0.9*	1.04*	0.76	0.94
Congo Basin	Factual	observed	0.01									
	SSP126	strong mitigation	0.01	0.01	0.01*	0.01*	0.01	0.01	0.92*	1.94*	1.02	1.42
	SSP370	middle-of-the-roadmedium-high	0.01	0.01	0.01*	0.01*	0.02	0.05	0.91*	1.37*	1.59	5.07
	SSP585	no mitigation	0.01	0.01	0.01*	0.04*	0.02	0.05	0.69*	3.85*	2.57	3.97

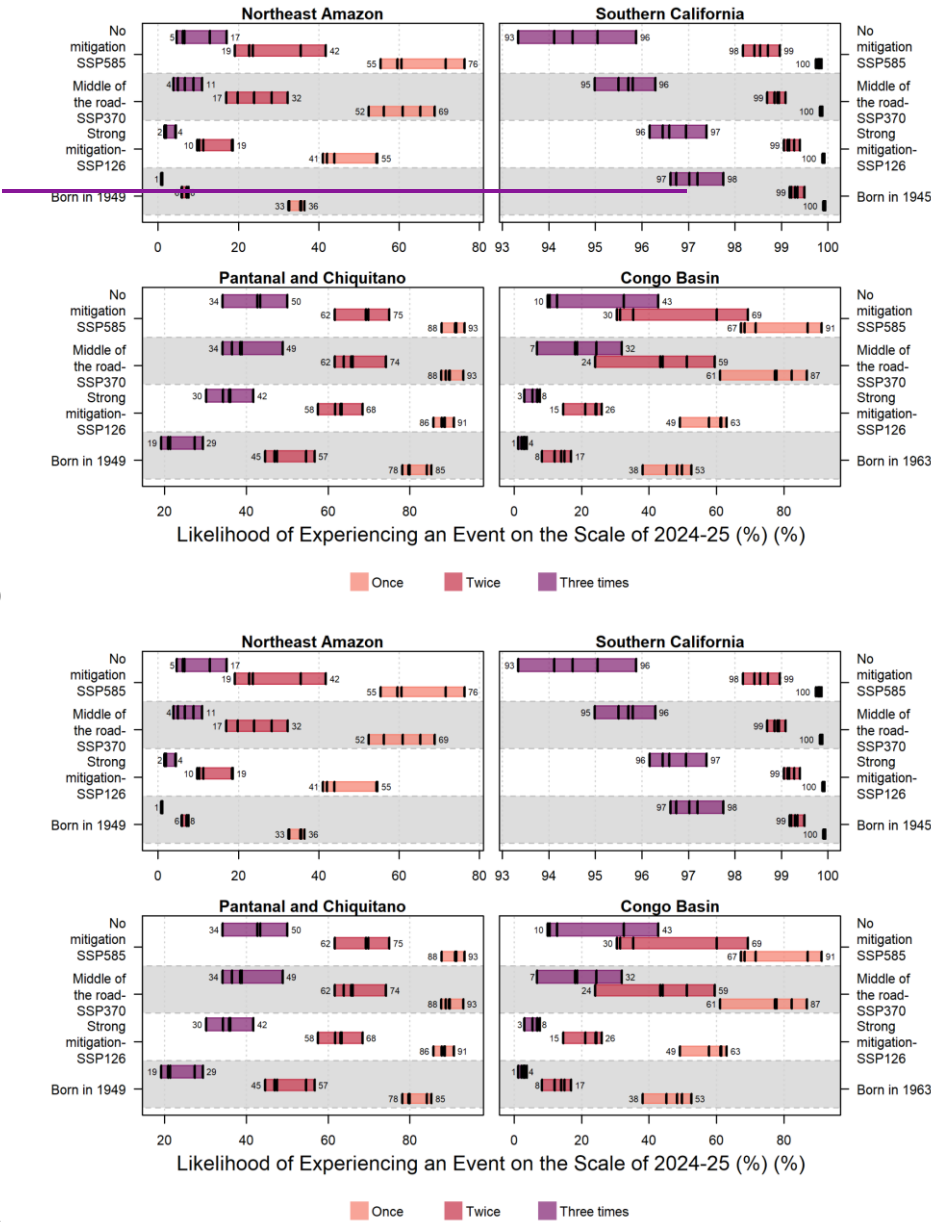


Figure 19: Likelihood of experiencing extreme fire events similar to those of 2024-2025 during the average lifetime of a citizen, based on current life expectancy (2023): Brazil (75.8 years, Northeast Amazonia, the Pantanal-Chiquitano), USA (79.3 years, Southern California), and Democratic Republic of Congo (61.9 years, Congo Basin). Bars show the probability of experiencing at least one, two, or three such events if born today under different scenarios:

historical climate (bottom bar in each group), SSP126, SSP370, and SSP585 (subsequent bars, bottom to top). Black vertical lines indicate individual GCM estimates; bar heights show the range across models.

6.2.3.2. *Pantanal and Chiquitano*

By mid-century (2050), no scenario shows significant increases in the frequency of BA levels such as 2024 at the regional scale (**Table 6**). All scenarios project modest increases by this point: about 1.14-1.15 times more frequent in SSP126 and SSP370, with slightly higher increases in SSP585 (up to 1.22 times). However, substantial changes emerge later in the century (**Figure 18**). Under SSP370, the likelihood of these fires becomes significantly higher by the 2070s, with a 1.2-fold (20%) increase relative to historical conditions. By 2100, SSP585 shows the greatest increases, up to 1.44 times more frequent, while SSP370 projects 1.34 times (**Table 6**). SSP126 demonstrates clear mitigation potential, limiting increases to about 1.13 times, with no significant change throughout the century.

For areas with the highest BA (top 5% grid cells), future changes in the frequency of 2024-like events are significantly different from 2010-2020 for both mid-century (2050) and by 2100 (**Table 6**). Increases at the sub-regional level are larger than regional averages, though not as dramatic as in Northeast Amazonia: by the end of the century events such as those from the 2024-25 season are expected to increase 1.26 to 1.75 times under SSP585, while SSP126 keeps increases much smaller (1.03-1.24 times). SSP370 projections fall between these (1.21-1.45 times), demonstrating that mitigation could still meaningfully limit the occurrence of extreme fire. Increases in the likelihood of extreme BA in high burning cells could begin as early as the 2030s under SSP126, driven in part by potential increases in fuel availability, though this effect could level off or reverse by mid-century (**Figure S29**). Under SSP370 and SSP585, increases in frequency of extreme BA start to take hold by the 2040s, though large changes may not emerge until after 2060.

These future extremes will mainly be driven by declining moisture availability over the entire region (**Figure 18**). For the most extreme BA areas, this moisture signal is less certain, and changes in fuel, though uncertain, could be large enough to modulate moisture effects (**Figure S29**).

Increases in BA will likely occur across the region by 2090 under all scenarios except in the wetland core of the Pantanal (**Figure S35**), where responses are much more uncertain. Areas of increased extreme fire behaviour exist even under SSP126, but most of the region is projected to see reductions or little change in extremes. In contrast, SSP370 drives widespread increases in extreme BA across almost the entire region, except the wetlands. However, as **Section 5** and other studies (e.g. Barbosa et al., 2022, 2025b) highlight, recent increases in extreme fire have been driven by the combined effects of climate change and wetland degradation, factor not considered in the future projections. This means increases in wetland fire extremes could arise sooner, even by the 2030s or 2040s under SSP126. Under SSP585, widespread increases in extreme BA may arise as soon as 2030 (**Figure S33**), and by 2040 even SSP126 shows large areas of the Pantanal and Chiquitano with much higher chances of a 1-in-100 event (**Figure S34**). Under the SSP126 scenario, the lower chances of extreme events by 2100 compared to mid-century (2040-2050) reinforce how strong mitigation strategies may alter wildfire trajectory throughout the 21st century in this region.

Finally, in terms of lived experience, someone born in the 1940s would already have had a high chance (78-85%) of witnessing a fire event like 2024 during their lifetime (**Figure 19**), with **Section 5** showing climate and human factors likely contributed substantially. Even under SSP126, this rises to 86-91% for someone born today. The difference is most striking for multiple-event likelihoods. Historically, someone born in the 1940s would have had a 19-29% chance of seeing three such events. Under SSP370, this rises sharply to 34-49%, similar to

SSP585 (34-50%). Even under SSP126, the likelihood of seeing two such events exceeds 50% (58-68%), compared to 45-57% historically.

6.2.3.3. Southern California

While January 2025 fire activity was likely influenced by anthropogenic climate change (Section 5.2.2.3), future projections suggest that similar-scale BA extremes may become less frequent (Table 6; Figure 18). However, this depends strongly on how local vegetation responds to rising CO₂ and climate change.

Looking ahead, models do not project a significant increase in the frequency of these regional-scale extremes (Figure 18). In fact, under SSP370 - a scenario closely aligned with current emissions trajectories, the likelihood of 2025-like events in terms of January BA slightly declines by a factor of 0.79 to 0.95 by the 2090s versus 2010s. Similar trends are seen under SSP585, though with the potential for stronger decreases. SSP126, however, showed no robust change by the end of the century.

The projected decline in extreme fire activity in Southern California appears to be driven primarily by modelled increases in tree cover, which occurs even with GCMs with declining precipitation, suggesting that it is largely driven by CO₂ fertilisation and enhanced water-use efficiency (Figure S28). This effect is more pronounced in drier climates like Southern California, where rising CO₂ concentrations reduce water stress on plants and promote vegetation growth. While this leads to greater fuel loads, our framework also represents tree covers influences on fuel moisture, which can suppress fire risk tipping the balance toward fewer extreme fire events in many model simulations. CO₂ concentrations are higher in SSP585 and SSP370 compared to SSP126, which explains why this effect is more pronounced in these scenarios. However, when tree cover is held constant at present-day levels, this signal weakens considerably. Under these “fixed tree” simulations, future projections of extreme fire activity become much more uncertain, with wide variation across scenarios all the way to the 2090s (Figure 18). Climate projections themselves for the region are mixed. Some models show increases in January precipitation and fewer dry days, while others suggest drier conditions (Figure S42). These divergent signals further contribute to uncertainty in fuel moisture and fire behaviour over the coming decade.

Our projections, therefore, rely on modelled tree and shrub cover from a global land surface model, which, while bias-corrected using historical observation (Figure S28), is primarily designed to capture broad-scale vegetation patterns. The model includes global plant functional types (PFTs) such as evergreen and deciduous shrubs, which encompass Mediterranean shrublands like those found in Southern California, but also represent structurally similar ecosystems in very different climatic and ecological settings (e.g., tropical savannas, tundra scrub). As a result, while the model tends to perform reasonably well in estimating total woody cover, it may not fully capture the fine-scale ecological gradients or the dominant shrubland dynamics that drive fire activity in this region. In particular, it may miss key features of chaparral systems and their interannual variability. Future work using regionally calibrated vegetation models or integrating remote sensing estimates of fuel structure may help increase confidence in projections for fire-prone shrub-dominated systems like Southern California.

Therefore, while our models suggest a potential future decrease in large-scale fire extremes in Southern California, this outcome depends on how burned area responds to increasing tree cover, and how vegetation itself responds to rising CO₂ and changing climate. Both relationships remain uncertain and will require further investigation. Understanding the evolving links between fuel load, fuel moisture, and ignition risk in the region is essential to refining future fire risk projections in this region.

6.2.3.4. Congo Basin

By the 2050s, none of the emission scenarios project a significant increase in the frequency of regional-scale 2024-like fire events (**Table 6**). Both SSP126 and SSP370 project modest changes, ranging from slight decreases to increases of up to 1.28 times more frequent, though wide uncertainty means small decreases remain possible. Substantial increases emerge by 2100, especially under higher-emissions scenarios. Under SSP370, the likelihood of large fire events rises by 1.11-1.52 times, with SSP585 showing similar values. In contrast, SSP126 holds the increase to just 1.03-1.11 times, indicating a meaningful mitigation potential.

For the most extreme fire events (top 5% of grid cells), projected increases in frequency are more substantial (**Table 6**). No scenario shows significant differences by 2050. However, significant and potentially large changes emerge by 2100. Under SSP370, the frequency of these high-BA extremes could rise by up to 5 times relative to historical conditions (range: 1.59-5.07), slightly higher than the 4-fold increase under SSP585 (2.57-3.97). SSP126 limits this increase substantially to just 1.02-1.42 times. These results show that even under a mitigation pathway, some increase in extreme BA is likely, but the scale of that increase is drastically reduced.

The primary driver of increased fire risk in the region is declining moisture availability, with drier conditions projected across much of the basin (**Figure 18**). In the higher-emissions scenarios (SSP370 and SSP585), increased fuel availability may amplify this effect. For the most extreme fire-prone areas, however, fuel controls show little change, suggesting that moisture stress will be the dominant factor shaping future fire behavior (**Figure S29**).

Spatially, increases in BA are relatively uniform across the region, though some local differences emerge (**Figures S39-41**). The eastern DRC may experience small decreases in July average BA, though increases remain more likely. In contrast, Gabon, Equatorial Guinea, and central DRC (particularly south of the Congo River) are projected to see the largest increases, with BA doubling or even quadrupling in some areas. Some of these increases, particularly along the Gabonese and Equatoguinean coasts, may begin as early as the 2030s.

In terms of lived experience, someone born in the DRC in 1963 with a life expectancy of 61.9 years, would have had a 38-53% chance of experiencing at least one event like that of July 2024 (**Figure 19**). For those born today, this rises to 49-63% under SSP126, 61-87% under SSP370, and as high as 67-91% under SSP585. The likelihood of experiencing multiple such events also increases markedly. Under SSP585, someone born today would have a 30-69% chance of seeing two events, and a 10-43% chance of seeing three. In contrast, SSP126 limits this to 15-26% for two events and just 3-8% for three, highlighting the powerful influence of mitigation. Indeed, the chance of seeing just one event under SSP126 is comparable to seeing two under SSP585.

7. Conclusions: Summary of the State of Wildfires in 2024-25

7.1. Extreme Wildfire Events of 2024-25

- **Global:** A total of 3.7 million km² burned globally during the 2024-25 fire season, 9% below the average of previous seasons (4.0 million km²), ranking 16th of all fire seasons since 2002. Despite the relatively low area burned, global fire carbon emissions were 2.2 Pg C, 9% above average and the 6th highest on record, driven by intense and high-emission fires in South America and Canada. This pattern reinforces a trend towards growing fire impacts in carbon-rich forest ecosystems, even during years with below-average fire extent globally.
- **South America:** South America experienced an unprecedented fire season setting a new record for carbon emissions. Emissions reached 263 Tg C (84% above average), with BA also 120,000 thousand km² (35%) above average. Bolivia, Brazil, and Venezuela each saw high or record-breaking anomalies, with Bolivia setting national records for both BA and C emissions. Record fire activity occurred across multiple biomes including the Chiquitano dry forests, Pantanal wetlands, and southern and Northeast Amazonia. These fires were characterised by extremely large, fast-spreading, and intense events despite fire counts often being average or below average, highlighting a pattern of fewer but larger and more intense fires on the continent. Highlights:
 - **Northeast Amazonia (Focal Event):** Record-breaking fire activity affected the moist tropical forests north of the Amazon River and Rio Negro, including large portions of Venezuela, Guyana, Suriname, and northern Brazil. Several ecoregions experienced all-time highs in burned area or carbon emissions, with fire activity peaking March-April and again in late 2024. Air quality impacts and environmental degradation were reported across the region.
 - **Pantanal-Chiquitano (Focal Event):** Extreme fire season across Bolivia and adjacent Brazil, with the Chiquitano dry forest and Pantanal wetlands (the world's largest wetlands) seeing some of the largest fires on record. Bolivia experienced the highest national carbon emissions total ever recorded (187 Tg C), with the Santa Cruz department (Bolivia) alone responsible for 157 Tg C. Fires destroyed critical habitat, caused severe air pollution, and threatened biodiversity hotspots. The pantanal recorded PM2.5 concentrations of 903.2 µg/m³ in September 2024, 60 times the WHO daily standard.
 - **Amazonas State, Brazil:** A record-breaking year for fire activity in this moist tropical forest region. Fire counts were up +154% versus the long-term average, and BA and fire size reached record levels. The 95th percentile fire size anomaly was +60%. This was one of the few regions in South America where high fire counts and severe individual fire behaviour co-occurred.
 - **Mato Grosso and Mato Grosso do Sul States, Brazil:** Both states saw record-breaking fire intensity and rate of spread. In Mato Grosso, 95th percentile fire size was 266% above average, despite fire counts being 54% below average. Mato Grosso do Sul experienced record emissions (+323%) and fire growth rates, pointing to fast, intense fires likely driven by land-use change and drought.
 - **Pará State, Brazil:** This state recorded its highest-ever BA (36,000 thousand km²) and major emissions anomalies (+61%). Fire activity expanded deep into forested areas, likely linked to land clearing. It was among the most significant subnational contributors to Brazil's fire totals in 2024-25.
 - **São Paulo State, Brazil:** Unusually high-intensity fires occurred despite a relatively small area burned. 95th percentile fire size and intensity both set new

records. Carbon emissions were nearly double the historical average (+190%), driven by a combination of unseasonal drought and land-use pressures.

- **Bolívar and Delta Amacuro, Venezuela:** Two states in northeast Venezuela experienced record emissions and BA, with Bolívar seeing a +133% BA anomaly and Delta Amacuro impacted by early-season fire peaks. These fires affected swamp forests and grassland regions.
- **Coastal and Andean Ecuador, Peru, and Colombia:** Subnational analysis reveals record or high-ranking anomalies in 8 provinces of Ecuador, 7 regions of Peru, and multiple Colombian ecoregions. These include areas in southwestern Amazonia and the eastern Andean slopes, where record fire sizes and intensities occurred despite average fire counts.
- **Guyana and Suriname:** Six ecoregions in Guyana and two districts in Suriname experienced record fire counts and BA, contributing to the focal Northeast Amazonia event but deserving standalone mention given the extent and duration of the anomalies.

- **North America:** The 2024-25 fire season was the second most severe on record for North America, with total C emissions of 194 Tg C (112% above average) and BA of 31,000 thousand km² (35% above average). Canada again saw extreme fire activity for the second year running, with 282 Tg C emitted and over 46,000 thousand km² burned, second only to the record-breaking 2023-24 season. In the US, the catastrophic Palisades and Eaton Fires in California in January 2025, which killed at least 3031 people, destroyed over 11,500,750 homes, and caused over \$140 billion in damages. Highlights:

- **Southern California, USA (Focal Event):** The most disastrous wildfire event in modern US history occurred in Los Angeles County in January 2025 during a severe Santa Ana wind event. The Palisades and Eaton Fires severely damaged or destroyed over 11,500,750 homes, killed at least 3031 people, displaced over 150,000, and caused economic losses exceeding US\$140 billion (including insured losses of US\$20-75 billion). Fires also severely degraded air quality, disrupted water supplies, worsened the housing crisis, and led to mass evacuations and air quality emergencies.
- **Western Canada:** Northwest Territories, British Columbia, Alberta and Saskatchewan experienced their second-highest emissions year on record with a combined emissions anomaly of +191 Tg C and provincial anomalies in the range of +184-441%
- **Mexico:** According to national statistics, Mexico experienced its worst wildfire season on record with over 8,000 wildfires and more than 16,500 km² burned. Particularly severe activity occurred in March-May, reportedly driven by drought and elevated temperatures. This record is not captured in our analyses based on global satellite products, warranting further investigation of the differences.
- **Alberta, Canada:** Extreme wildfires in summer 2024 destroyed 358 structures and led to \$1.23 billion in damages, second only to the Fort McMurray fire of 2016. The town of Jasper was evacuated. Two firefighter fatalities occurred.
- **New York, USA:** In an unusual late-season outbreak, every borough experienced multiple wildfires during a two-week span in October-November 2024, an unprecedented fire signal in a densely populated urban environment.

- **Africa:** For the second consecutive year, fire extent in Africa was well below average, with BA in the African savannah biome 12% below average, the third lowest on record. However, several regions experienced notable fire anomalies, particularly the Congo Basin, northern Angola, and South Africa. Record-setting BA and C emissions were recorded in some regions of the Republic of Congo and the Democratic Republic of Congo. Despite the extent of these events, many went under-reported in the media, reinforcing the importance of Earth observation-based monitoring. Highlights:

- **Congo Basin (Focal Event):** Record fire activity and C emissions in the Republic of Congo and Democratic Republic of the Congo. Fires contributed to the region's highest primary forest loss since 2015 and caused hazardous air pollution, with DRC reporting PM_{2.5} levels 11 times the WHO standards. Fires in western ecoregions such as the Atlantic Equatorial and Central Congolian lowland forests were particularly intense.
 - **South Africa:** Fires killed 34 people, including 6 firefighters, and destroyed thousands of livestock and homes. KwaZulu-Natal Province was particularly affected. High fuel loads from previous wet years reportedly contributed to the intensity.
 - **Côte d'Ivoire:** Fires in Séguéla (Worodougou region) burned ~~50,000 ha~~500 km², destroyed homes and plantations, and killed 23 people. Other fatal incidents occurred in Bouna, Bongouanou, and Taabo.
- **Asia:** Overall, Asia experienced a below-average fire season, with BA 26% below average and C emissions 28% below average. However, significant regional extremes were observed. Highlights:
 - **Nepal:** Nepal endured its second-worst fire season since 2002, with over 1,000 wildfires. Wildfires killed more than 100 people, with significant destruction of forests and homes. In the Lumbini Province, wildfires devastated ~~11,448 ha~~114 km² of forests and destroyed more than 230 houses and livestock shelters.
 - **Northern India:** Uttar Pradesh experienced its most severe wildfire season on record, reportedly driven by crop burning, heatwaves, and dry fuel accumulation. Regional fires contributed to severe haze episodes in New Delhi in November 2024, with PM_{2.5} concentrations exceeding 200 µg/m³ across large parts of Northern India (13 times the WHO daily standard).
 - **Iran:** Worst fire season since 2002. Fires burned key national parks and forest areas. Carbon emissions, fire counts, and BA all reached record highs, reportedly driven by a combination of climate stress and human pressures.
 - **South Korea and Japan:** Japan's largest wildfire in over 50 years took place in Iwate Prefecture in February 2025, destroying 221 buildings. South Korea's deadliest wildfires occurred in March 2025 (just outside of the 2024-25 fire season), killing 31 and damaging 4,000 homes.
 - **Sichuan and Guizhou, China:** A fire in Sichuan lasted 14 days, displaced 3,000 people, and impacted multiple villages. Strong winds and dry spring conditions reportedly drove unusually large wildfires.
 - **Heilongjiang and Jilin, China:** Record BA occurred in both provinces. Though not widely reported, these events underscore rising fire activity in northeast Asia, which has been linked to agricultural burning and shifting policy enforcement.
 - **Republic of Sakha and Zabaikalsky krai, Russia:** Fires in these regions accounted for 65% of total forest area burned across Russia and forced 58 redeployments of firefighting resources involving 1,861 firefighters.
- **Europe:** Europe recorded its fourth lowest BA since 2002, with 30,000 km² burned (49% below average) and C emissions 22% below average. However, there were stark regional contrasts. Highlights:
 - **Portugal:** Most destructive fire season since 2017. Over ~~137,000 ha~~1,370 km² burned, with 16 fatalities and €180 million in damages. Fires in September affected wildland-urban interface areas in the northwest. A ~~5,000 ha~~50 km² fire in Madeira entered the laurel forest, a rare cloud forest and UNESCO World Heritage site. This incident highlighted the vulnerability of non-fire-adapted ecosystems under increasing fire pressure.
 - **Serbia, North Macedonia, and Bulgaria:** Worst wildfire seasons in two decades. Large-scale fires led to EUCPM activations and widespread

evacuations, including four fires >1 thousand km² (>10,000 ha) in North Macedonia alone.

- **Ukraine:** Nearly 1 million ha 10 thousand km² burned during 2024-25, mostly in conflict-affected eastern areas. Fires were likely exacerbated by warfare, with higher-than-usual forest losses reported.
- **Romanian Danube Delta:** An unusually dry winter led to 45,000 ha 450 km² of wetlands burning in February 2025. Though a recurring phenomenon, this was one of the most extensive burn events yet, and emblematic of changing fire regimes in sensitive wetland ecosystems.
- **Turkey Türkiye (Mardin Province):** A rapidly spreading fire in June 2024 burned farmland and villages, killing 15 people and injuring at least 70. It was one of the deadliest fire events in the Eastern Mediterranean this season.
- **Austria and Germany:** While Central Europe had a quiet fire year overall, Austria recorded its highest number of fires and largest BA since 2012, and Germany had a slightly above-average season, consistent with a slow but steady upward trend.

- **Oceania:** Oceania experienced a moderate fire season overall, but numerous high-impact events were recorded. Highlights:
 - **Western Australia:** Over 1,000 large fires burned ~470,000 ha amid record heat and severe dryness between Perth and Carnarvon. The Skeleton Rocks fire (44,000 ha) impacted long fire-interval ecosystems and a lithium mine, while the largest fire near Cervantes burned 80,000 ha and disrupted regional honey production. Manjimup fires affected over 42,000 ha of native forest and required interstate response.
 - **Central Australia:** Over 5.7 million ha burned by October 2024, including a 450,000 ha fire near Devil's Marbles that forced closures of major infrastructure. In January, 80,000 ha burned in the West MacDonnell Ranges, including national parks and Aboriginal land trusts.
 - **Victoria and Tasmania:** Severe dry lightning outbreaks triggered major fires in culturally sensitive landscapes. Victoria's Grampians National Park saw two-thirds of its area burned, and the Little Desert fire burned 90,000 ha in under 8 hours. Tasmania's northwest fires burned 100,000 ha, affecting the Tarkine and Cradle Mountain.
 - **Queensland:** Firefighters responded to 40 incidents at Mount Isa, with one fire burning over 100,000 ha for nearly two months. Smoke exposure caused hospital admissions and endangered species such as the Carpentarian Grasswren were threatened.
 - **New Zealand:** Peat fires at Whangamarino Wetland and Tiwai Peninsula each burned ~1,000 ha, likely generating significant CO₂ emissions after similar events in 2022 emitted 0.6 million tonnes CO₂.

7.2. Focal Regions

In this year's report, our detailed analyses target three tropical regions and Southern California. The extreme nature of events in these focal regions are given in Section 2..

- **Northeast Amazonia** saw record forest fire activity, with burned area +332% above average, the highest since records began. Fires severely impacted Indigenous communities, displacing thousands and degrading air and water access.
- The **Pantanal-Chiquitano** experienced its worst fire season on record, with burned areas nearly triple the average and carbon emissions six times above average. Fires affected both the Pantanal wetlands, the world's largest tropical wetland, and the Chiquitano dry forests of Bolivia. PM_{2.5} pollution reached hazardous levels of 900 µg/m³, carrying strong potential for detrimental health and economic impacts.

- **Southern California** recorded catastrophic wildfire losses, with 3031 deaths, 11,599,750 homes destroyed, and US\$140 billion in total damages. PM_{2.5} levels peaked at 483 µg/m³, triggering a regional housing and insurance crisis.
- The **Congo Basin** had its highest recorded fire activity at 28% above the annual mean, contributing to a +150% increase in primary forest loss in 2024 versus 2023. Fires were the main driver of deforestation but received minimal media or institutional attention, highlighting a broader lack of media coverage of fires affecting equatorial Africa.

7.3. Impact Assessments

In this year's report, we incorporate new assessments of the impact of fires on society, specifically via the exposure of populations, physical assets, and carbon projects to fire and via smoke degrading air quality. Key findings from our analyses were as follows.

Population exposure:

- We estimate that ~100 million people were exposed to wildfire activity globally during the 2024-25 fire season, with the highest exposures in India and the Democratic Republic of the Congo (15 million each).
 - Uttar Pradesh (India) recorded the highest subnational exposure at 4.6 million people, a 146% increase over average, followed by Heilongjiang (China, 3.7 million) and Punjab (India, 3.6 million).
 - Despite severe fire seasons, Canada, Brazil, and Bolivia contributed modestly to global population exposure due to the remoteness of areas burned.
 - Other countries experiencing large *relative* anomalies in population exposure included: Jordan, Peru and Ecuador (Andes); Venezuela, Guyana, and Suriname (northern South America), Nepal and Niger.
- 20 thousand,000 people were officially displaced according to IDMC displacement records, or 0.02% of those exposed according to our analysis. This reflects a gap between exposure and formal displacement, though true disruption is likely higher than in the IDMC records due to known issues with underreporting.
- Exposed communities may still suffer serious health, economic, and psychological consequences (e.g., missed income, increased debt, long-term health declines), even if they are not formally displaced.

Physical asset exposure:

- According to our analysis, an estimated US\$215 billion in physical assets were exposed to wildfires in 2024-25. Top countries by asset exposure were India (US\$44 billion), United States (US\$26 billion), China (US\$17 billion), South Africa (US\$14 billion).
 - Other countries with high absolute asset exposure were: Mexico, TurkeyTürkiye, and Russia (~US\$8 billion each).
 - Other countries experiencing large *relative* anomalies in physical asset exposure were: Pakistan, Sudan, Chad, Albania, Greece, Iraq, Syria, and Eritrea.
- US\$57 billion in direct losses were recorded in the international disaster database EM-DAT, including \$53 billion caused by fires affecting LA and southern California.
 - Direct financial losses are generally smaller than our estimates of physical asset exposure (the detection of fire in proximity to the built environment) because exposure is a measure of potential for loss, and not of loss itself.
 - In the case of southern California, recorded direct financial losses from fires were three times larger than our estimates of exposed physical assets due to the underestimation of asset density in our analysis. A lesson from this work is that analyses of exposure must account for the significant variation in the

4058 density of real estate value across states of the USA, and likely in other
4059 countries as well.

4060 Carbon project exposure:

- 4061 • The 2024 fire season saw record BA across forestry projects in the Voluntary Carbon
4062 Market (VCM): 169 of 927 projects (18%) experienced fire, the highest on record since
4063 2001, with burned area in 2024 affecting 1.6% of project areas on average.
- 4064 • 72% of projects experienced above-average drought contributing to elevated risk of
4065 fire, with 13% exceeding extreme drought thresholds (SPEI < -2).
- 4066 • The 2024 fire season had an above average impact on carbon projects in Latin and
4067 northern America, average BA was recorded in Eurasia and below average in Africa.
4068 In addition to climate, land use and land cover changes and project activities also
4069 contributed to regional differences in observed extremes.
- 4070 • Despite elevated BA in the latest fire season, 46% of all carbon projects experienced
4071 no fire in the entire period since 2001, while 67% experienced little fire (defined as
4072 <0.5% burned annually in the surrounding 50-km buffer).
- 4073 • The 2024 season underscores that while high-integrity forest carbon projects remain
4074 a key climate change mitigation tool, the permanence of carbon stored or avoided is
4075 increasingly threatened by extreme fire years, especially under worsening climate
4076 extremes.

4077 Air quality:

- 4078 • Our analysis of air quality impact in this report focuses exclusively on the Pantanal-
4079 Chiquitano focal region, where population-weighted PM_{2.5} exceeded the WHO daily
4080 standard (15 µg/m³) on 43 days between July to October (over a third of all days in the
4081 period) from July to October and peaked at a regional population-weighted average of
4082 61 µg/m³ in September, with fires accounting for ~58% of the pollution. Smoke
4083 emissions from fires were the sole cause of exceedances of the WHO daily standard
4084 on 50 days in the period July-October.
- 4085 • Wildfire smoke emissions exposed communities to extremely harmful air quality in
4086 various world regions, according to direct measurements (**Appendix A**). For example,
4087 communities in the Brazilian Pantanal, Southern California, Bolivia, and northern India
4088 were exposed to PM_{2.5} concentrations of over 60, 30, 30, and 13 times the WHO daily
4089 standard of 15 µg/m³, respectively.

4091 7.4. Diagnosing Causes and Assessing Predictability

- 4092 • **Weather was the dominant driver of fire activity during all of the 2024-25 focal**
4093 **events** targeted in this report, contributing 40% to 70% of the explainable cause.
 - 4094 ○ **Fuel** availability and dryness increased in importance during the most severe
4095 fires (up to 40% of explainability) and determined the final extent of BA.
 - 4096 ○ **Ignitions** were consistently dominated by human influence, but they did not
4097 emerge as a primary cause of fire activity during the 2024-25 focal events (only
4098 around 10% of explainability).
- 4099 • In **Northeast Amazonia** fire activity was predominantly driven by persistent, large-
4100 scale drought conditions that depleted deep soil moisture reserves. These droughts
4101 suppressed fuel moisture recovery for extended periods, even during rain periods. Soil
4102 moisture anomalies reached up to 3 standard deviations below the climatological
4103 mean, with values dropping to as little as 2% of average. The prolonged drought
4104 significantly increased fuel dryness and, during the period of most intense burning, fuel
4105 importance rose up to 20% above its annual baseline. Fuel also determined the final
4106 burned area extent, contributing significantly to the observed anomalies in BA,
4107 accounting for up to 50% in the sub-regions where BA anomalies were most extreme.
4108

Human-caused ignitions were present but did not emerge as a leading cause of fire (10-20%). Their contribution remained limited and at times negative compared to what is considered usual (thus reducing the total extent), likely reflecting limited ignition opportunities or active suppression efforts to limit BA.

- In the **Pantanal-Chiquitano**, extreme fire activity was primarily driven by antecedent drought persisting since 2023. Deep soil moisture remained in the driest 15% of records, 1-2 standard deviations below average, despite wetter conditions in early 2024. Although February-April rains moistened surface fuels, they failed to recharge deeper layers. Weather dominated fire activity (71% average contribution), with fuel importance rising to 40% during the peak burning week in early August and explaining over 50% of final BA anomalies. Lightning played a minimal ignition role, often occurring in association with convective downpour. Human-caused ignitions, though still dominant, were lower than in previous years and at times limited burned area extent.
- In the **Congo Basin** extreme fire activity was driven by prolonged and severe drought persisting over recent years. The usual spring wet season (March-May) failed to occur, and the second wet season later in the year provided limited relief, leaving deep soil moisture up to 2 standard deviations below climatological norms. Weather was the dominant driver of fire activity, with rainfall 67% below and temperatures 90% above climatological averages, placing vegetation and soil dryness among the driest 1-2% of records (2003-2023). Human activity accounted for most fire ignitions but as for the other 2 tropical regions they were not the main causes of the fire severity and actually acted to reduce the final BA
- In **Southern California**, the 2024-25 fire season was marked by atypical seasonality, with extreme fire activity occurring in January well outside the usual summer peak. The Palisades and Eaton fires were driven by a rare convergence of weather, fuel, and ignition factors, each contributing significantly (weather: 40%, fuel: 30%, ignition: 20%). Despite preceding years of exceptional wetness, a short-lived but extreme drying of surface fuels (3 standard deviation deviations below normal) and intense winds (3 standard deviation deviations above normal) created highly flammable conditions. These fires ignited and spread rapidly at the wildland-urban interface, highlighting how brief windows of extreme weather can override generally moist background conditions and trigger major off-season events in these parts of the world.
- There were **distinct challenges to the forecasting of all focal events**:
 - In **Northeast Amazonia**, our models correctly identified two high-risk fire seasons, but most of the burning occurred during the first (February-April), not the second (August-November), despite similar fire danger forecasts. This disconnect highlights a key limitation: high fire danger does not always lead to high fire activity. Human factors, such as suppression, fire bans, or shifts in land use, likely played a role and are currently underrepresented in fire prediction systems.
 - In the **Pantanal and Chiquitano**, fires were closely linked to long-term drought conditions that dried out fuels months before the fire season peaked. Fire activity rose only after this slow build-up, meaning accurate forecasts required capturing both drought and fuel dynamics. While the general heightening of fire danger was picked up by the FWI, the machine-learning-based PoF model, which includes fuel conditions, better predicted when and where fires would actually occur.
 - In **Southern California**, fire prediction remains difficult without accounting for the 'whiplash effect' that arises from extreme fire weather following on from wet periods with high vegetation productivity. A wet period led to vegetation growth, followed by rapid drying and strong winds that enabled the January fires. As in the Pantanal-Chiquitano, including fuel information helped the PoF model identify higher-risk areas more accurately than the FWI.

- In the **Congo Basin**, both FWI and PoF tended to overpredict fire danger. While drought increased flammability, ignition remained limited, possibly due to cultural practices, suppression efforts, or fewer ignition sources (though reporting on such activities in this region is extremely limited). Here, human activity and fuel moisture, more than fire weather, shaped outcomes. The FWI system, which unlike PoF does not include these factors, was less effective in predicting fire activity in the Congo Basin.

7.5. Attribution to Global Change

- **Climate change has increased the likelihood of extreme fire events across all focal regions studied.** The high fire weather and extreme levels of burning seen in 2024-25 were significantly more likely in a world with human-induced climate change.
- In **Northeast Amazonia**, we ~~find that~~~~found that~~ the extreme fire weather during January-March 2024 was 30-70 times more likely due to anthropogenic climate forcing, while the risk of regional BA totals being as observed in the period was 2.1 times greater due to anthropogenic climate forcing and the area burned by fires was four times greater.
 - Our attribution analysis shows high confidence that climate change played a major role in Northeast Amazonia's record fire season. We are virtually certain (>99%) that anthropogenic climate forcing increased the risk of extreme fire weather, very likely (96%) that it amplified the area affected, and likely (89%) that it increased the chance of the extreme burned area observed.
 - While climate change has clearly enhanced the probability of extreme events in the region, such as that seen in 2024, there was conversely no robust evidence that climate change increased average annual BA totals in Northeast Amazonia during 2003-2019.
 - An increase in annual average BA during 2003–2019 of up to 17% was attributed to ~~socioeconomic~~~~changes in land use~~ changes since 1900-1917, indicating that long-term human activities have elevated typical fire levels in the region.
 - Overall, our attribution analyses suggest that climate change has enhanced the likelihood of extreme fire events in the region, against a backdrop of increased annual BA levels driven by socioeconomic change such as land use/land cover change and human ignitions.
- In the **Pantanal and Chiquitano**, we find that the extreme fire weather August-September 2024 was 4-5 times more likely due to anthropogenic climate forcing, while the risk of regional BA totals being as observed in the period was 3.3 times greater due to anthropogenic climate forcing and the area burned by fires was around 34 times greater.
 - Our attribution of extreme fire weather to climate change ~~was made~~~~virtually~~~~was~~~~virtually~~ certain (>99%, IPCC definition), while the amplification of both extreme burned area and region-wide burned area extent was attributed with likely confidence (87%). Taken together, these findings provide strong evidence that anthropogenic climate change raised the odds of the largest fire season on record in the Pantanal-Chiquitano region.
 - In addition to the enhanced odds of extreme BA events, a 10% increase in annual average BA during 2003-2019 was attributed to climate change.

- At least a two-fold increase in BA during years with 2024-like fire conditions was attributed to socioeconomic [change variables](#), indicating that human activities [may](#) have substantially increased the risk of widespread fire under extreme conditions. However, other analyses focusing on long-term annual average burned area suggest that some human-driven changes [may could](#) have reduced typical annual fire activity. While these findings are not strictly contradictory since they examine different aspects of the fire regime, the contrast between them reduces confidence in attributing overall fire trends to socioeconomic drivers alone and points to the need for further investigation. [Improved confidence in socioeconomic attribution for the Pantanal-Chiquitano will require higher-resolution data on human activity, seasonal and temporal variation in land use and fire practices, and explicit representation of water management and wetland dynamics, in combination with validation against observed extreme events](#)
 - Overall, extreme BA events in the Pantanal-Chiquitano, such as those seen in August-September 2024, are made more likely by climate change and are superimposed on broader background increases in fire extent related to climate change and possibly socioeconomic changes in the region.
- In **Southern California**, we find that the risk of regional BA totals being as observed during January 2025 was 2.3 times greater due to anthropogenic climate change and the area burned by fires was 25 times greater.
 - Our attributions of amplified BA extent during the event to climate change were all made with at least 89% confidence. It is therefore *likely* (per IPCC definitions) that anthropogenic climate change raised the odds of the costly wildfires in Southern California during January 2025.
 - The meteorological conditions during the event were previously studied by the World Weather Attribution (WWA) group, who reported that extreme fire weather conditions were also made more likely, by around 40%, with other indicators such as prolonged drought and delayed seasonal drying also showing climate influence (Barnes et al., 2025). We did not perform an independent attribution of fire weather here due to a lack of data required for construction of a counterfactual scenario in our attribution protocol.
 - In addition to the enhanced odds of extreme BA events, a 7% increase in annual average BA during 2003-2019 was attributed to climate change.
 - Our BA attribution approaches did not provide robust evidence that socioeconomic change affected average annual BA, though this is possibly due to the difference between the coarse model resolution and the fine scale over which effects would be expected at the wildland-urban interface in this region.
 - Overall, extreme BA events in Southern California, such as those seen in January 2025, are made more likely by climate change and are superimposed on broader background increases in fire extent related to climate change.
- In the **Congo Basin**, we find that the extreme fire weather July-August 2024 was 3-8 times more likely due to anthropogenic climate change, while the risk of regional BA totals being as observed in the period was 60% greater due to anthropogenic climate change and the area burned by fires was three times greater.
 - It is virtually certain that anthropogenic climate change contributed to

the extreme fire weather observed during the 2024 season in the Congo Basin. The widespread extent of burned area was very likely influenced by climate change (92% likelihood), while the most extreme sub-regional burned area events were likely influenced (78% likelihood). Together, these findings indicate that climate change increased the odds of the largest fire season on record in the region.

- In addition to the enhanced odds of extreme BA events, a more than 45% increase in annual average BA during 2003-2019 was attributed to climate change.
- Our BA attribution approaches did not provide robust evidence that socioeconomic change affected average annual BA during 2003-2019 versus a pre-industrial counterfactual. [Our BA attribution approaches did not provide robust evidence that socioeconomic change affected average annual BA during 2003-2019 versus a pre-industrial counterfactual.](#)
- Overall, extreme burned area events in the Congo Basin, such as those seen in July–August 2024, are made more likely by climate change and are superimposed on broader background increases in fire extent attributable to climate change, with no robust evidence that socioeconomic changes significantly altered recent fire activity.

7.6. Seasonal and Multi-Decadal Outlook

- **Fire weather and BA forecasts for boreal summer 2025 highlight several areas with elevated probability of anomalous fire danger.** Probabilities for anomalous fire prone seasons are high across Canada, northeast Europe (including the UK), and parts of Siberia. These conditions following the second-warmest May on record globally (1.4 °C above pre-industrial levels), with exceptional dryness and the lowest northwestern European rainfall since 1871.
 - In equatorial regions, forecasts show a more than 60% chance of anomalous fire weather conditions in Northeast Amazonia, the Congo Basin, and the Himalayan foothills.
 - In the US, severe drought conditions in Arizona and Texas are already leading to elevated fire activity in line with predicted anomalies in fire weather.
 - Seasonal outlooks of burned area anomalies coincide with fire weather anomalies in western South America, southern California, central North America, and eastern Central Asia.
 - Chile and northern Australia stand out with >50% confidence for anomalous fire activity during the boreal summer of 2025.
 - Despite high FWI in central Europe, we could not confidently predict a BA anomaly due to insufficient historical fire-climate data for reliable modelling.
- **In Northeast Amazonia, our climate model projections consistently indicate a rise in extreme wildfire risk by the end of the century.** Under a ~~middle-of-the-road~~[medium-high](#) emissions pathway (SSP370), the frequency of regional BA totals on the scale of 2024 are projected to increase by up to 57% by 2100.
 - Also under SSP370, the greatest rate of increase (factor of 2-3 rise) is projected in the sub-regions that burned most extensively in the extreme event of 2024 (5% of model cells with greatest BA).
 - Under a *no mitigation* scenario (SSP585), an even sharper rise is projected, with a near-doubling of the frequency of extreme (2024-like) events at the regional scale. Greater rates of increase (up to a four-fold rise) are projected in the sub-regions that burned most extensively in 2024.

- In contrast, limiting warming under a *strong mitigation* scenario (SSP126) effectively contains future fire risk. By 2100, the increased frequency of an extreme (2024-like) event is limited to 9%, with the sub-regions that burned most extensively in 2024 showing no significant change. This demonstrates the strong potential of climate action to mitigate the risk of future extreme fires in Northeast Amazonia.
 - Projections of future increased risks are not spatially uniform in any scenario. In some areas, such as Amapá and northern Pará in Brazil, and southern Suriname, increased extreme fire activity is projected as early as the 2030s under higher-emissions scenarios (SSP370 and SSP585). Even under SSP126, rises in extreme BA are projected for parts of the moist forest zone.
 - The frequency of extreme (2024-like) events is projected to rise only modestly in all scenarios through 2050, however by 2100 the increased risk under higher emissions scenarios (SSP370 and SSP585) clearly emerges from that of SSP126.
- In the **Pantanal and Chiquitano, our climate model projections indicate further increases in extreme wildfire risk by the end of the century.** Under a *middle-of-the-road/medium-high* emissions pathway (SSP370), the frequency of regional BA totals on the scale of 2024 are projected to increase by up to 34% by 2100.
 - Also under SSP370, the greatest rate of increase (21-45% rise) is projected in the sub-regions that burned most extensively in the extreme event of 2024 (5% of model cells with greatest BA).
 - Under a *no mitigation* scenario (SSP585), an even sharper rise is projected, with a 44% rise in the frequency of extreme (2024-like) events at the regional scale. Greater rates of increase (up to a 75% rise) are projected in the sub-regions that burned most extensively in 2024.
 - In contrast, limiting warming under a *strong mitigation* scenario (SSP126) effectively contains future fire risk. By 2100, the increased frequency of an extreme (2024-like) event is limited to 13% and is not significant, while the sub-regions that burned most extensively in 2024 experience minimal increases in frequency (up to 24% rise). This demonstrates the strong potential of climate action to mitigate the risk of future extreme fires in the Pantanal-Chiquitano.
 - At the regional scale, only modest increases in the frequency of extreme (2024-like) fire seasons are projected by mid-century across all scenarios. However, by 2100, the increased risk becomes more pronounced under higher emissions pathways, with clear divergence between scenarios.
 - At the sub-regional level in the areas that burned most extensively, earlier increases in extreme fire risk could begin as soon as 2030.
 - Projections of future increased risks are not spatially uniform in any scenario. Geographically, widespread increases in BA are projected across most of the Pantanal-Chiquitano by 2100, though the response is considerably more uncertain in the Pantanal than in the Chiquitano. Some areas of increased extreme (2024-like) fire frequency may still emerge in the Pantanal even under SSP126.
 - It is important to note that these projections do not fully incorporate local *in situ* drivers, such as wetland degradation, which have already contributed to more frequent fires in recent years. Increases in fire activity might be expected to occur earlier than the models indicate, especially along the wetlands and adjacent drainage areas.
- In **Southern California, our climate model projections of future change in extreme (2024-like) fire events are highly uncertain.**
 - While high-emissions simulations under SSP585 and SSP370 suggest that extreme fire events could become less frequent over time, this strongly

depends on how vegetation responds to rising CO₂ and a changing climate.

- In particular, simulations suggest that increased tree cover driven by CO₂ fertilisation under higher-emissions scenarios (SSP585 and SSP370) may raise fuel loads while simultaneously increasing fuel moisture, with the overall effect being to reduce the likelihood of extreme fire events in our models.
- However, when removing changes in tree cover, the projected future frequencies of extreme (2024-like) events become highly uncertain with no consistent direction of change under future scenarios.
- There is a critical need for improved observation and modelling of how vegetation structure, fuel moisture, and local ecological processes shape fire behaviour in Southern California. Nonetheless, Southern California remains highly exposed to fire risk. Even under scenarios that suggest a decline in fire extremes, most residents alive today are still likely to experience multiple extreme fire seasons like 2025 in their lifetime. Stronger local adaptation and more regionally tailored research on climate-vegetation-fire interactions will be essential to manage risk in the coming decades.

- In the **Congo Basin, our climate model projections indicate that further increases in extreme wildfire risk are likely by the end of the century.** Under a ~~middle-of-the-road~~ medium-high emissions pathway (SSP370), the frequency of regional BA totals on the scale of 2024 are projected to increase by up to 50% by 2100.

- Also under SSP370, far greater rates of increase (up to a 5-fold rise) are projected in the sub-regions that burned most extensively in the extreme event of 2024 (5% of model cells with greatest BA).
- Projections under SSP370 and SSP585 show similar levels of elevated risk, indicating that mitigation efforts stronger than those implied by SSP370 are likely needed to meaningfully reduce future fire risk.
- In contrast, limiting warming under a *strong mitigation* scenario (SSP126) effectively contains future fire risk. By 2100, the increased frequency of an extreme (2024-like) event is limited to at most 11%, while the sub-regions that burned most extensively in 2024 experience a far smaller increase in frequency (up to 42% rise) than under higher emissions scenarios. This demonstrates the strong potential of climate action to mitigate the risk of future extreme fires in the Congo Basin.
- Projections of future increased risks are not spatially uniform in any scenario. Some of the largest projected increases, seen in Gabon, Equatorial Guinea, and central DRC, may begin as early as the 2030s, with the frequency of extreme (2024-like) events is projected to increase 2 to 4-fold by 2100. This increase is driven largely by declining fuel moisture as climate change reduces rainfall and increases dry spells across much of the region.

- **Anthropogenic climate change has the potential to significantly increase future fire risk for living generations, turning previously exceptional events into events that are experienced several times in a generation.**

- **Northeast Amazonia:** Our projections ~~show~~ indicate that a person born in this region ~~today has during the 1940s had~~ a ~~41-55-33-36%~~ likelihood of experiencing at least one extreme fire episode on the scale of January-March 2024 ~~in their lifetime~~. For someone born today, this likelihood rises to 41-55% under *strong mitigation* scenario (SSP126). This likelihood rises to 52-69% under a ~~middle-of-the-road~~ medium-high scenario (SSP370 - closest tested to our current emissions pathway), and 55-76% under a *no mitigation* scenario (SSP585). The odds of experiencing two or more such events are considerably higher under *no mitigation* (19-42%) than under *strong mitigation* (10-19%).
- **Pantanal-Chiquitano:** Our projections indicate that a person born in this region during the 1940s already had a ~78-85% likelihood of experiencing at

least one ~~extreme fire season-like~~ episode on the scale of August-September 2024. For someone born today, this likelihood rises to 86-91% even under SSP126. Under SSP370, the likelihood of experiencing at least two 2024-scale fire seasons rises to 62-74%, compared to 45-57% for someone born in the 1940s, but even under low emissions, the chance of two such events exceeds 58%. These findings highlight that while climate change mitigation can reduce future fire risk, it is not sufficient on its own. Early adaptation, ecosystem management, and stronger fire governance will be essential to reduce future impacts.

- **Congo Basin:** Our projections ~~show~~ indicate that a person born in this region ~~today has during the 1940s had~~ a ~~49-63-38-53%~~ likelihood of experiencing at least one ~~extreme fire season-like~~ episode on the scale of July 2024-. ~~For someone born today, this likelihood rises to 49-63%~~ under SSP126. This likelihood rises to 61-87% under SSP370 and 67-91% under SSP585. The likelihood of experiencing multiple events differs starkly across SSPs, with up to a likelihood of 43% for three events under SSP585, compared to just 3-8% under SSP126.

7.7. Progress in the State of Wildfires Report

This report incorporates a number of major advances in our annual reporting on the State of Wildfires in the prior fire season. In **Section 2**, we added a new analysis of fire intensity to our extreme event identification variables, and we evaluated the dependence of our extreme event identification on the source of BA observation by incorporating data for 2019-2025 from two additional BA datasets (FireCCIS311 and VIRS VNP64A1), supplementing our consistent multi-decade analysis based on the MODIS BA dataset (MCD64A1). The contribution of regional expert knowledge was also formally recognised through the establishment of regional expert panels for each continent, with these panels consulted for their interpretation of results across all aspects of the report. We added **Section 3**, which presents an entirely new set of impact assessments relating to population exposure, asset exposure, carbon project exposure and air quality degradation. In **Section 4**, we expanded the analysis of the predictability of the focal event to include seasonal predictions of burned area, complementing the fire danger seasonal forecasts already provided. In **Section 5**, our main advancement was a new approach to attributing both extreme regional BA totals and sub-regional BA extremes directly to the 2024-25 focal events, made possible by developing near real-time counterfactuals and employing methodologies for aggregating probabilities across space. This represents a step-change versus our first report, which focussed on attributing sub-regional BA extremes only and substituted near real-time counterfactuals with less targeted counterfactuals for the 2003-2019 period. By creating more robust counterfactuals with observed events, and accounting for the stochastic nature of fire anomalies locations, we were able to more directly and confidently assess whether human influence made these specific fires more likely on regional scales. In **Section 6**, we extended our forward-looking capabilities by providing seasonal forecasts of BA, complementing the fire danger forecasts already presented in previous reports. We also added future projections of FWI at future global warming levels of 1.5-4.0°C, providing a clearer picture of how extreme wildfire risk may evolve in the coming decades.

This new report documents the progress made in the observation, diagnosis, modelling and attribution of extreme wildfire events and their impacts. As a community, our work is both driving innovation in the methods under use and prompting the development of new capabilities for the routine analysis of extreme wildfire events and their impacts. This new report builds on our inaugural report (Jones et al., 2024b) and documents the progress being made by the fire science community.

By combining cutting edge techniques in fire forecasting, prediction and modelling across the sections of our report, we compile multiple lines of evidence for a clear climate signal in recent

fire extremes. Our complementary analyses consistently point to a strong role for climate change in driving extreme fire conditions, showing that human influence, both through climate change through socioeconomic change factors, are increasing fire risk and producing extreme wildfires. While individual methods sometimes diverge, particularly in regions like the Pantanal, where local socioeconomic factors emerge more clearly as drivers in some analyses, the overall convergence across independent lines of evidence builds confidence in the conclusion that climate changes exerts significant upwards pressure on the likelihood of extreme fire events.

These multiple lines of evidence show that human influence, often through climate change though sometimes through socioeconomic factors, are increasing fire risk and driving higher burned areas. Across every region analysed, we find clear signals that recent extreme wildfires are not purely natural events, but increasingly shaped by human-driven changes to climate and ecosystems.

A key strength of this report lies in its systematic evaluation of model performance across diverse regions of the globe. In this edition, for instance, we identify limitations in the capacity of coarse-resolution air quality models to assess smoke exposure in small regions (**Section 3**), and show how projections of future fire activity can be strongly influenced by how models represent sensitive vegetation responses to uncertain climate changes (**Section 6**). In regions such as California, long-term projections are particularly sensitive to changes in tree cover, which can be affected by uncertainties in both climate inputs and modelled vegetation responses.

A rich body of observations, such as land surface and meteorological data, are available to observe and model the effects of climatic change and variability on extreme fire likelihood, in particular following important advances in the modelling of fuel load and moisture dynamics during recent decades. However, a major outstanding barrier that consistently limits the effectiveness of our analyses, and those of the broader fire science community, is a severe lack of information regarding *in situ* human activities. Funding of [research](#) projects that overcome this barrier is paramount and carries the greatest potential to drive a step-change in performance of fire models and predictive systems. Often, prediction and modelling analyses rely on basic indicators of human effects such as population density, which cannot sufficiently represent the diversity of relationships between people, their land uses, and the outcomes for fire ignitions and spread dynamics. Our work, and that of many others, highlights the need to develop global datasets that effectively represent the range of human-fire interactions that occur on Earth but with sufficient scalability to support regional and global analyses. Inevitably, there will be a trade-off between the geographical scalability and nuance of these datasets as they are developed.

Overall, our international collaboration ~~routine~~[routinely](#) catalogues fire extremes and annually evaluates the most extreme fire events of international relevance using state-of-the-art fire science tools. We provide a consistent stream of actionable information to policymakers, disaster management services, firefighting agencies, and land managers, informing action on enhancing society's resilience to wildfires through investment in preparedness, mitigation, and adaptation.

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● Appendix A: Year in Review by Continent

This appendix includes the review completed by regional expert panels to supplement our quantitative analyses of extremes in the 2024-25 fire season (**Section 2**). Details of the assembled panel are provided in **Table A1**, below.

Table A1: Experts contributing to the identification of extreme events and characterisation of the global fire season during March 2024-February 2025.

Region	Co-authoring Experts	Country of Organisation / Nationality	Professional Background(s)	Supporting Expert Panellists	Others Consulted
Africa	Kebonye Dintwe	Botswana		Lucy Amisshah (Ghana), Sally Archibald (South Africa), Natasha Ribeiro (Mozambique), Tercia Strydom (South Africa)	
	Aya Brigitte N'Dri	Ivory Coast			
Asia	Cong Gao	China		Bambang Saharjo (Indonesia), Sundar Sharma (Nepal), Raman Sukumar (India), Veerachai Tanpipat (Thailand), Bo Zheng (China)	
	Elena Kukavskaya	Russia	Research		
Europe	Paulo Fernandes	Portugal	Research	Davide Ascoli (Italy), Stefan Doerr (UK), Julien Ruffault (France), Gavriil Xanthopoulos (Greece)	
	Cristina Santín	Spain	Research		
	Johan Sjöström	Sweden	Research		
North America	Crystal Kolden	USA	Research, Firefighting	Jacqueline Shuman (USA), Matt Jolly (USA), Piyush Jain (Canada), Chelene Hanes (Canada)	
	Mathieu Boubonnais	Canada			
	Victoria Donovan	USA			
Oceania	Hamish Clarke	Australia	Research, Environmental Management		Simeon Telfer, South Australian Country Fire Service; Rui Feix, Western Australian Department of Fire and Emergency Services; Chris Collins, Tasmania Fire Service; Grant Pearce, Fire and Emergency New Zealand; David Field, New South Wales Rural Fire Service; Russell Stephens Peacock, Queensland Fire Queensland and Fire and Emergency Services; Maggie Towers, Northern Territory Police, Fire and Emergency Services
	Sarah Harris	Australia	Research, Emergency Management		
South America	Liana Anderson	Brazil	Research	Dolors Armenteras (Colombia), Francisco de la Barrera (Chile), Mauro Gonzalez (Chile), Celso H.L. Silva-Junior (Brasil)	
	Carlos M. Di Bella	Argentina	Agronomist/Research		
	Bibiana Bilbao	Colombia Venezuela	Research		
	Galia Selaya	Bolivia	Tropical Ecology/Research and action		

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4545 ○ A1. Africa

4546 National and regional fire monitoring statistics are rarely recorded or made publicly available
4547 by fire agencies in Africa, meaning that our assessment of the latest global fire season largely
4548 focuses on the insights provided by global data analyses. According to these data, the total
4549 BA in Africa was approximately 2.4 million km² during the 2024-2025 fire season, 11.6% below
4550 the mean annual BA since 2002 (2.7 million km²). Most of the BA occurred in non-forest (2
4551 million km²), with the remaining portion in the forest. Non-forest and forest BAs were 12% and
4552 7% lower than the mean annual BA, respectively. The BA anomaly was notably larger in NHA
4553 (-14.6%) than in SHA (-9.1%). The relatively low BA in many parts of the continent could be a
4554 result of a combination of factors, though it aligns with a trend that has been attributed to the
4555 continued suppression of fire from expanding croplands (Andela et al., 2017) and to changing
4556 rainfall patterns across the continent (Zubkova et al., 2019).

4557
4558 Africa's most pronounced positive anomalies in BA and fire C emissions of the 2024-25 fire
4559 season were seen in the Congo basin and northern parts of Angola (**Figure 1, Figure 2; Table**
4560 **2, Table 3**). BA in the Republic of Congo was 25% above average, the highest on record, and
4561 similarly fire C emissions were 25% above average (**Table 2, Table 3, Figure S43**). In the
4562 Democratic Republic of the Congo, the Mai-Ndombe and Sankuru provinces each
4563 experienced record levels of BA or fire C emissions with anomalies in the range of 36-58%
4564 (**Table 2, Table 3**). These anomalies were centred on several western ecoregions of the
4565 Congo Basin, including the Atlantic Equatorial coastal forests where BA was more than triple
4566 the annual mean, Western Congolian swamp forests where BA was twice the annual average
4567 and the Central Congolian lowland forests where BA was 77% above average, and the
4568 Northwestern Congolian lowland forests where BA was 55% above average. These results
4569 align with the recent report of the Global Forest Watch (World Resources Institute, 2025) which
4570 found that the Democratic Republic of the Congo (DRC) and the Republic of the Congo
4571 experienced their highest rates of primary forest loss since 2015. While loss to wildfire is a
4572 minor component of overall forest loss in the region (below 15%), for instance compared to
4573 the expansion of shifting cultivation, wildfires were the major explanation for the more than
4574 doubling (+150%) increase in forest loss in 2024 versus 2023.

4575
4576 The uptick in fires in the Congo basin can be linked in part to the enabling effect of record-
4577 breaking fire weather caused by drought in the region (**Section 2.2.2.1**), however a range of
4578 socioeconomic changes have also been underway and likely influenced the events of last
4579 year. Use and degradation of the forests for resources, often linked to an increase in related
4580 wildfire ignitions, is growing due to the extraction of timber to produce charcoal, clearing of
4581 land for the expansion of cash crops, and shortening or cessation of fallow periods in
4582 smallholder shifting cultivation systems (World Resources Institute, 2025). The potential
4583 effects of fires in this region on forest carbon stocks are globally significant (though they are
4584 yet to be quantified), with the region's swamp forests harbouring 30 billion metric tonnes of C
4585 in peat (Garcin et al., 2023). The 2024 IQAir World Air Quality Report highlighted that the
4586 Democratic Republic of the Congo had an annual average PM_{2.5} concentration of 58.2 µg/m³,
4587 over 11 times higher than the World Health Organization's annual standard (IQAir, 2025). This
4588 indicates hazardous air quality levels, in part to the effects of elevated wildfire smoke
4589 emissions (IQAir, 2025). Despite the potentially large impacts on society and the environment,
4590 there was very limited news coverage on the impacts of these fires by national news outlets
4591 across the region. This underscores the importance of projects such as ours and the Global
4592 Forest Watch (World Resources Institute, 2025) using Earth Observations to routinely trace
4593 environmental extremes and highlighting events that would otherwise have gone under-
4594 reported.

4595
4596 The high BA in the Congo Basin during 2024-25 has implications for various initiatives
4597 supported by non-governmental organisations in the region, which aim to promote protection
4598 and sustainable management of tropical forests. For example, UNEP's Congo Basin

Sustainable Landscapes Programme (Green Policy Platform, 2025) supports action in Cameroon, Central African Republic (CAR), the Democratic Republic of the Congo (DRC), Equatorial Guinea, Gabon, and Republic of the Congo. In programmes such as this, wildfire can sometimes be considered a secondary disturbance factor compared to other factors such as clear-cut deforestation, but years such as 2024 demonstrate that large-scale intermittent fires in the region can have lasting consequences for forest loss.

In Angola, BA and fire C emissions were 15-49% above average in the provinces of Moxico, Huila, and Bié and were either record-setting or high-ranking (**Figure S44; Figure 2, Figure 3; Table 2**). As discussed in **Section 2.2.2** and investigated formally for the Congo Basin in **Section 4**, a particularly hot and dry fire season elevated the risk of fire in these regions and coincided with broader social and economic factors promoting fire ignitions. The poor economic situation in Angola over the past three years has prompted deregulation of the charcoal industry and the harvesting of trees for charcoal production has risen, driving up fire activity (Valor Económico, 2024; VisiteHuila, 2024). In addition, the government has been promoting agriculture through financial programs, leading to the clearing of land through shifting agriculture in Miombo woodlands (Fundo de Garantia de Crédito, 2024; World Bank, 2024b). In certain provinces, particularly Moxico in Angola, burning for hunting purposes is also widespread and declining populations of prey have been linked to increased burning of areas that were previously hunted less regularly (Papelo, 2024). These are just some of the socioeconomic factors that may have contributed to the elevated availability of ignition sources during 2024-25 fire season, when fire weather was particularly conducive to fire.

In Algeria, fires have killed and injured dozens and caused significant loss of life and damage in recent years. At least 34 people were killed and several hundred were injured in Bejaia province during the 2023-24 fire season (Jones et al., 2024b). However, in 2024-25, a low number of fires were recorded and there were no casualties in Algeria, which could be attributed to various factors such as the availability of better firefighting equipment, new fire management policies, and a new law that was passed that imposes life imprisonment for those caught deliberately starting forest fires (Serrah, 2024; The Arab Weekly, 2024). Algerian authorities launched a wildfire prevention system that included 13 water-bombing aircraft and 100 drones for monitoring and tracking firefighting operations. For instance, 26 fires were extinguished within 24 hours in the central and eastern regions of Algeria, with no injuries or casualties reported (Gabriel, 2024).

In South Africa, the total BA was over 46,000km², which was 17% higher than the mean annual BA. According to a report by the organisation Working on Fire (2024), the increased intensity and frequency of these fires continue to challenge firefighting resources. The 2024-25 fire season broke records, with 2,750 firefighting teams dispatched, with a record number of 34 people losing their lives, including firefighters. In KwaZulu-Natal Province, the wildfires claimed the lives of 14 people, of whom 6 were firefighters who were trapped in a blaze. In addition to the lost lives, thousands of people were displaced, over 2,050 livestock destroyed, and critical infrastructure damaged. The high intensity fires in South Africa could be due to a string of particularly high rainfall years that resulted in large accumulated grassy fuel loads.

In Côte d'Ivoire, the overall BA in 2024-2025 was lower than the historical average, contrary to what some national experts had expected following the long dry season which began earlier than usual in the savanna areas of the country where fire is most widespread (N'Dri et al., 2018, 2024; Soro et al., 2021). Nonetheless, the fire season was marked by an above-average fire size distribution and there were several deadly events in the country's main fire hotspots, with fires burning over ~~150,000~~ **1.5 thousand km²** of forest, ~~2,800~~ **280 km²** of plantations, ~~109~~ **1 km²** of reforestation projects, and 107 properties in 2025 (CNDFB, 2025). In the department of Séguéla (Worodougou region), wildfires in February 2024 destroyed ~~50,000~~ **500 km²** of natural vegetation, ~~261~~ **3 km³** of cropland, ~~236~~ **2 km²** of cashew plantations, 19 homes, and claimed the lives of 23 individuals across the villages of Sélakoro, Djénigbé,

Touna, Djoman and Kondogo. In Bouna (Bouunkani region), fires affected around ~~12,387~~ ~~ha~~120 km², of which ~~7,528~~ ~~ha~~75 km² were forested, leading to additional humanitarian impacts. Three further fatalities were recorded between February and March 2024 in Bongouanou (Moronou region) and Taabo (Agnéby-Tiassa region). These impacts occurred despite the continued efforts of the *Comité National de Défense de la Forêt et de lutte contre les feux de Brousse* (CNDFB), such as the construction of firebreaks during the dry season and awareness campaigns. This reflects the challenges posed by expanding agricultural land and ignition sources, fire suppression policies that focus on fire exclusion to protect valuable crops (e.g. cashew nuts) but promote fuel build-up, and a lack of prescribed burning in Côte d'Ivoire's savanna ecosystems (Ruf et al., 2010; Soro et al., 2020; Kouassi et al., 2022). Generally, fire activity and BA have been declining across all ecoregions of Côte d'Ivoire, which has been attributed to conversion of savannas to agricultural lands and also bush encroachment in savanna areas (N'Dri et al., 2022; Douffi et al., 2021; Kouassi et al., 2022).

○ A2. Asia

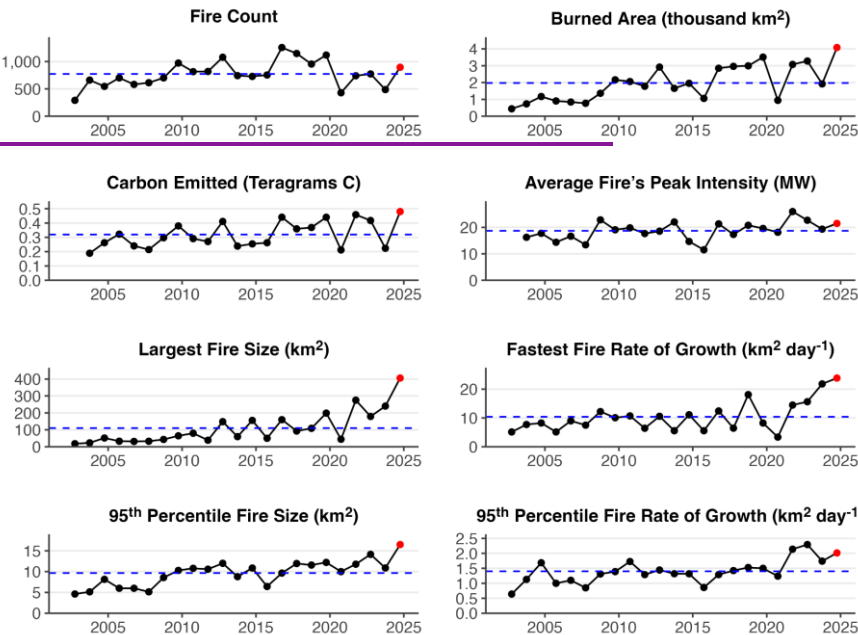
The 2024-25 fire season in Asia was generally not an extreme one, with much of Asia experiencing typical or low BA. Nonetheless, there were regional extreme fire events in the fire season.

Iran emerged as a notable case, experiencing its most severe wildfire season since 2002, marked by record-breaking BA, number of fires, and carbon emissions at the national level (**Figure 2, Figure 3**). Ecologically sensitive regions were disproportionately affected, including Karkheh National Park in Khuzestan Province and the forests and rangelands of Ab Kenar and Khan Ahmad Basht in Kohgiluyeh and Boyer-Ahmad Province (Global Fire Monitoring Center, 2024). As one of the driest countries in the world, Iran experiences approximately 1,500 wildfire outbreaks annually, resulting in the burning of ~~15,000~~ ~~ha~~150 km² of forest (Kheshti, 2020; Tavakoli Hafshejani et al., 2022). Human activities are the primary driver of wildfires nationwide, with deforestation, illegal logging, and accidental ignition contributing to the high incidence of fires (Masoudian et al., 2025). These anthropogenic pressures are compounded by systemic shortcomings against wildfires, including inadequate resource allocation and insufficient prevention measures, which challenge the protection of natural ecosystems (Iran International, 2024).

Nepal also endured its second-worst fire season since 2002 (**Figure S45**), with over 5,000 fires according to some sources (Bolakhe, 2024; and >1,000 individual fires in our analysis, **Figure S45**) causing more than 100 fatalities (Bolakhe, 2024). In Lumbini province, located in western Nepal, wildfires devastated ~~11,448~~ ~~ha~~114 km² of forests and destroyed more than 230 houses and livestock shelters (Sanju Paudel, The Kathmandu Post, 2024). These catastrophic events were driven by extreme drought, prolonged heatwaves, and frequent lightning strikes (Karuna shechen, 2024). Concurrently, anthropogenic factors including agricultural residue burning, poachers' use of fires, and unintentional human negligence, have exacerbated wildfire occurrence (Shradha Khadka, Governance Monitoring Centre Nepal, 2024). Notably, Nepal's forest cover has doubled over the last three decades, increasing from 26% to 45% between 1992 and 2016 (Karan Deep Singh, The New York Times, 2022). While Nepal's afforestation initiatives represent a significant environmental achievement, addressing the escalating human-nature conflict and strengthening resilience to climate-induced disasters remain critical challenges for ensuring the sustainability of this fragile progress.

Northern India, bordering Nepal, also experienced extreme heatwave and drought in 2024, triggering unprecedented wildfire activity across several states (Reuters, 2024). Uttar Pradesh, for example, experienced its most severe wildfire season, marked by record-breaking BA, carbon emissions, rate of growth, and fire size (**Figure A1**). Human activities, mainly land clearing and negligence, serve as the primary ignition source in Northern India,

4709 leading to uncontrolled wildfires. These fires are further exacerbated by the accumulation of
4710 dry pine needles in forests, which act as a ready fuel, and the steep Himalayan slopes, which
4711 accelerate the rate of fire growth (Vivek Saini, Climate Fact Checks, 2024). Agricultural
4712 practices in Northern India, a critical crop-producing region, have also contributed to the
4713 extreme wildfire season. Despite regulatory bans, post-harvest burning of crop residue has
4714 continued unabated in recent years (Arshad R. Zargar, CBS News, 2024). At the same time,
4715 temperature inversions coupled with Himalayan topographical blockages have trapped
4716 pollutants over Northern India. This phenomenon culminated in severe air haze episodes in
4717 New Delhi in November 2024, with PM_{2.5} concentrations exceeding 200 µg/m³ across large
4718 parts of Northern India (CAMS, 2024).
4719



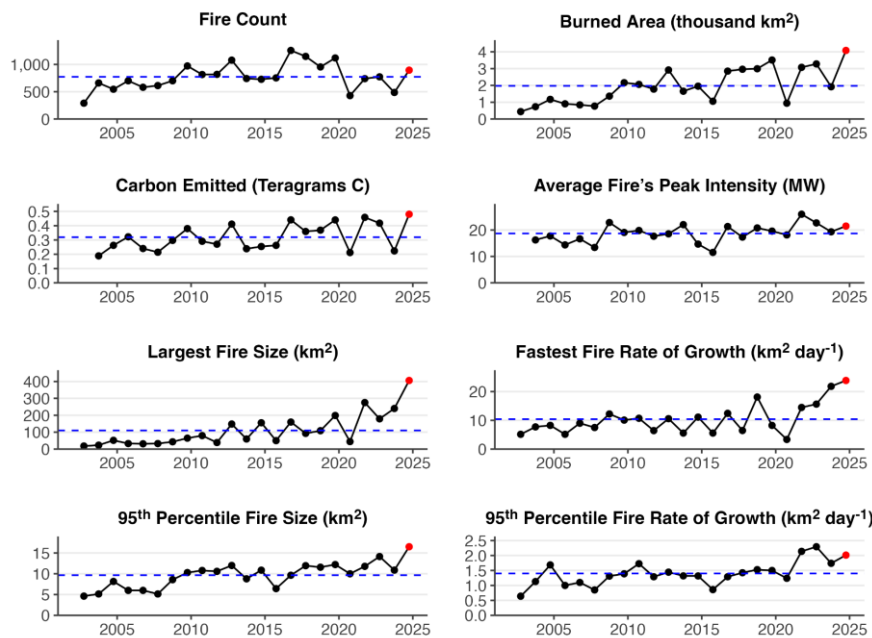


Figure A1: Summary of the 2024-2025 fire season in the Indian State of Uttar Pradesh. Time series show annual fire count, BA, C emissions totals within the region, as well as the average fire's peak fire intensity (95th percentile value of fire radiative power within fire perimeters), the 95th percentile fire size, fastest daily rate of growth, and 95th percentile fire daily rate of growth. Black dots show annual values prior to the latest fire season, red dots the values during the latest fire season, and blue dashed lines the average values across all fire seasons. Although Russia generally experienced a typical fire season, several regions in Siberia recorded extreme fire activity. Two regions (Republic of Sakha and Zabaikalsky krai) accounted for 65% of the total forest area burned in Russia (Avialesookhrana, 2024) with 97% of the fires recorded in hard-to-reach areas according to official data from the Federal Forestry Agency (Rosleshoz, 2024). The high fire activity was associated with intense heat, decreasing precipitation and dry thunderstorms (Rosleshoz, 2024), which have become more frequent phenomena in Siberia in recent decades (Huang et al., 2024). Firefighting was complicated by strong winds and mountainous terrain (Rosleshoz, 2024). To attract additional fire-fighting forces, federal emergency regimes were introduced from May 31 to November 8 in the Zabaikalsky krai and from June 28 to September 13 in the more northern Republic of Sakha, including in the Arctic Circle. In total, 139 redeployments involving 3,500 firefighters were carried out in 2024. The main causes of forest fires were lightning (48%), local population (39%) and fire transitions from other land categories (10%) (Rosleshoz, 2024). While the area burned in 2024 in the Republic of Sakha was not the highest compared to fire activity in the previous years, there is an increasing trend of fire activity and severity in the region over the last decade (ISDM, 2024), associated with weather anomalies (Tomshin and Solovyev, 2022) resulting in an increase in the duration of the fire season and the average area burned (Kirillina et al., 2020; Narita et al., 2021). The estimated total emissions for June 2024 were the third highest in the past two decades, following those of 2019 and 2020 (AMAP, 2024). In the Zabaikalsky krai, the total area burned in 2024 was about 7% of the area of the region, which is the highest value since 2010 (ISDM, 2025). Overall, both regions are considered hotspot areas of fire-induced change, where anthropogenic patterns and climate change are

4750 increasing ecosystem damage from wildfires and inhibiting recovery of natural ecosystems
4751 (Kukavskaya et al., 2016; Burrell et al., 2022).

4752
4753 Persistent dry and warm spring conditions in southwest China, particularly in Sichuan and
4754 Guizhou provinces, resulted in high-ranking BA anomalies (**Figure 2**). Strong winds
4755 exacerbated fire risks by increasing regional fire size and rate of spread, leading to large and
4756 fast-moving wildfires (Global Times, 2024). One of the most severe wildfires in Sichuan lasted
4757 14 days, displacing more than 3,000 civilians across 11 villages and one community (Dou et
4758 al., 2024). Northeast China, including Heilongjiang and Jilin provinces, also experienced
4759 widespread wildfire anomalies during the spring season (**Figure 2; Table 2**). Contrary to the
4760 climate-driven wildfires in southwest China, these wildfires were predominantly anthropogenic
4761 originating from crop residue burning. The Chinese government implemented policies in 2013
4762 and 2018 to control straw burning, a major contributor to air pollution, which initially achieved
4763 measurable success (Huang et al., 2021; Song et al., 2024). However, due to financial strain
4764 on rural communities and administrative pressures on local officials, recent policy adjustments
4765 have shifted from a zero-tolerance approach to a more flexible framework. This revised
4766 strategy permits controlled crop residue burning in designated areas during periods of low air
4767 pollution risk (Ding, Sixth Tone, 2025).

4768
4769 Earth observations data showed high-ranking BA anomalies, frequent fires, fires with large
4770 sizes, and rates of growth during 2024-25 in several regions of Lebanon, Palestine, Jordan,
4771 Iraq, Syria, United Arab Emirates, Philippines, and Laos (**Figure 2, Figure 4**), consistent with
4772 reports of persistent heatwave in these regions (Zachariah et al., 2024).

4773
4774 A drought that persisted from the 2024-25 fire season to the 2025-26 fire season has resulted
4775 in several highly impactful events in Asia (Faranda et al., 2025). Thus, from March 21st 2025,
4776 South Korea experienced its deadliest wildfires on record with very strong wind, burning
4777 across 11 regions and resulting in 31 deaths, 44 injuries, more than ~~3.3 thousand~~300
4778 displaced people, and at least ~~4 thousand~~000 homes damaged (Yonhap, 2025). The wildfire
4779 in Iwate Prefecture, Japan, which started on February 26th 2025, was the country's largest
4780 wildfire in over 50 years, killing one person, destroying 221 buildings and forcing evacuation
4781 of over ~~4,5 thousand people~~500people (NHK, 2025). These events are not reviewed at length
4782 here, however they will be featured in future editions of the State of Wildfires report.

4783 4784 o **A3. Europe**

4785
4786 In 2024, wildfire activity in the European Union was close to the long-term average in terms of
4787 total BA, but characterized by strong regional contrasts; approximately ~~420,000 ha~~4,200 km²
4788 were burned, slightly above the 18-year average (San-Miguel-Ayanz et al., 2025), with some
4789 countries experiencing record-breaking seasons and others seeing minimal fire activity.
4790 Across the continent, including in ~~Turkey~~Türkiye and Ukraine, a total of ~~1.82 million ha~~18
4791 ~~thousand km²~~ burned from March 2024 to February 2025 as recorded by the European Forest
4792 Fire Information System (2025), of which 48% pertain to large (~~>500 ha~~5km²) fires. The EU
4793 Civil Protection Mechanism (EUCPM) was activated 16 times in response to wildfires,
4794 providing international assistance to Greece, Portugal, Cyprus, Bulgaria, Albania and North
4795 Macedonia (European Commission Emergency Response Coordination Centre, 2024).

4796
4797 The 2024 wildfire season in the Nordic and Baltic countries was the calmest in recent decades.
4798 While spring was drier and warmer than average in some areas, abundant summer
4799 precipitation limited fire spread. No major wildfire events were reported, and most incidents
4800 were confined to small wildfires caused by land-management activities (Swedish Firefighters,
4801 2024). Likewise, wildfire activity in Western Europe during 2024 and early 2025 was subdued,
4802 as precipitation during spring and summer limited fire occurrence and spread across the
4803 region. France experienced one of its quietest seasons in recent decades, and similar

conditions were observed in Belgium, the Netherlands, Ireland, and the UK (Global Wildfire Information System, 2025). The fire season was insignificant in Central Europe, because of wetter-than-average conditions during spring, especially in the Czech Republic and in Slovakia. However, Austria saw the highest number of fires and the largest BA since 2012 (Global Wildfire Information System, 2025) and Germany experienced a slightly above-average fire season, consistent with the trend of the previous five years. The most notable incident was a wildfire in Harz National Park in July, which led to the evacuation of approximately 500 people and involved 150 firefighters (Deutsche Welle, 2024).

In Southern Europe fire activity varied widely depending on seasonal precipitation and fire weather, with notable peaks in July-August (Balkans) and September (Portugal). In Portugal (Figure S46), 2024 was the most impactful year since 2017: ~~137,111 ha~~ ~~1,400 thousand km²~~ burned on the mainland, around 20% above the past decade's average, with 25 fires exceeding ~~1,000 ha~~ ~~10 km²~~, eight of which surpassed ~~5,000 ha~~ ~~50 km²~~ (Instituto de Conservação da Natureza e Florestas, 2024). Most of these fires occurred as a sudden burst in mid-September in the northwest and under exceptional fire weather conditions (Instituto Português do Mar e Atmosfera, 2024). The Sever do Vouga complex and other major fires affected wildland-urban interface areas, resulting in 16 fatalities (Agência para a Gestão Integrada de Fogos Rurais, 2025), and €180 million in estimated losses across housing, infrastructure, forestry, and agriculture (Centro de Coordenação Regional Centro, 2024; Centro Pinus, 2024). Additionally, ~~48,272 ha~~ ~~480 km²~~ of protected areas and Natura 2000 habitats burned (Gonçalves and Marcos, 2024). In Madeira island, a fire burned over ~~5,000 ha~~ ~~50 km²~~, entering the non fire-adapted laurel forest, a UNESCO World Heritage Site (Público, 2024).

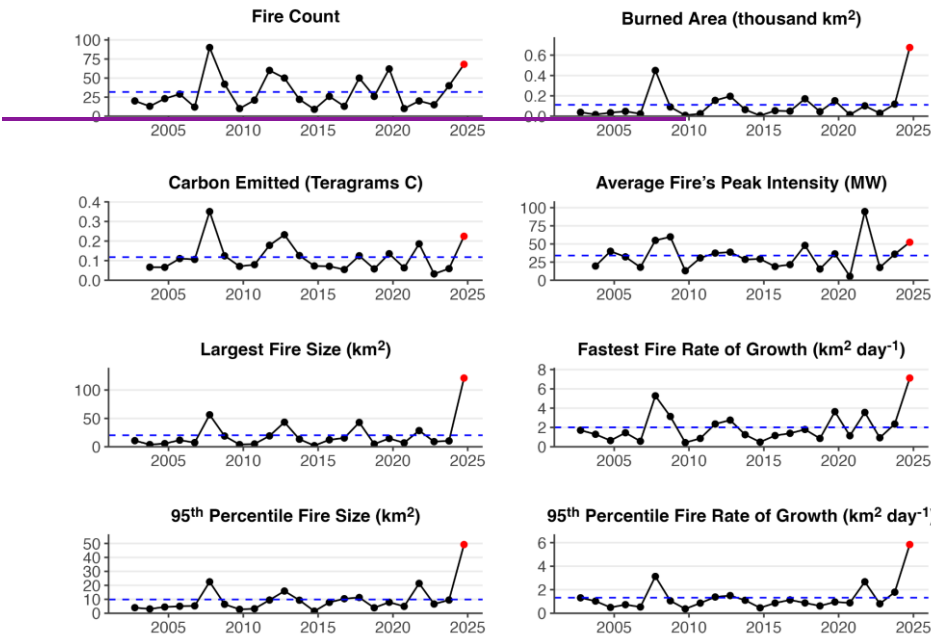
BA in Spain, Italy and Greece was respectively 41, 51 and 73% of the 2012-2023 average (Global Wildfire Information System, 2025). In Greece, the drought lasted until mid November, lengthening the fire season and enabling unusual high-elevation fires in the north. Nonetheless, strengthened preparedness and fire suppression hindered the spread of many potentially large fires. The most destructive fire occurred near Varnavas in August, entering the NE suburbs of Athens and killing one person (Giannaros et al., 2024).

The 2024 fire season in the Balkans and Southeastern Europe was among the most severe in recent decades for several countries. Wildfire activity was substantial in North Macedonia, Serbia, Albania, Kosovo and Bulgaria, including multiple large-scale events requiring international firefighting assistance. In Albania, the largest wildfire surpassed ~~4,000 ha~~ ~~40 km²~~ in the Dropull i Poshtëm region and the EU Civil Protection Mechanism was activated in response, with aerial support from Greece and Italy (Directorate-General for European Civil Protection and Humanitarian Aid Operations, 2024). Evacuations were carried out near the coastal town of Shengjin. Bulgaria experienced its worst fire season since 2007, with two wildfires in July destroying houses, the Sakar Mountain fire (Radio Bulgaria, 2024) and the Gorska Polyana fire (Novinite, 2024). North Macedonia (Figure A2) and Serbia faced their worst fire seasons in over two decades, and a state of emergency was declared in the former (Euronews, 2024), where four fires larger than ~~100 km²~~ (10,000 ha) were recorded (European Forest Fire Information System, 2025). On 16 August, Serbian authorities reported 135 active wildfires within 24 hours (N1info, 2024). Other countries in the region, such as Croatia and Montenegro, had seasons closer to the norm. In the Romanian Danube delta, and during an unusually dry winter, ~~45,000 ha~~ ~~450 km²~~ of wetlands burned in February 2025, a recurring phenomenon with increasing extent (Volodymyr and Andiy, 2025).

BA in ~~Turkey~~ ~~Türkiye~~ reached ~~270,000 ha~~ ~~2,700 km²~~, about 65% of the previous 12-years average (Global Wildfire Information System, 2025) but with noticeable societal consequences. Most large fires (up to ~~7000 ha~~ ~~70 km²~~) occurred in the province of Mardin (European Forest Fire Information System, 2025), including a rapidly spreading fire that burned farmland and impacted villages on 20 June, killing 15 and additionally injuring at least

70 people (The Nation, 2024). A fire that started near the coastal city of İzmir on 15 August brought havoc to the wildland-urban interface and ended up burning houses and injuring 78 people (Ozerkan, 2024).

Long periods of high fire danger combined with intensified aggression and scarcity of firefighting resources set the scene in Ukraine. The fire season was severe in extent and nearly 1 million ha 10 thousand km² burned between March 2024 and February 2025 (European Forest Fire Information System, 2025). This is larger than the combined BA in all of Europe, Middle East and North Africa (San-Miguel-Ayanz et al., 2025). As the majority of these fires are located near the front line in the eastern part of the country, warfare was presumably a major driver of their ignition, with forests seemingly accounting for a larger share of BA than in the recent past (The Guardian, 2025). Nonetheless, higher BA had been recorded in the past, namely >2 million ha 20 thousand km² in both 2014 and 2015 (Global Wildfire Information System, 2025).



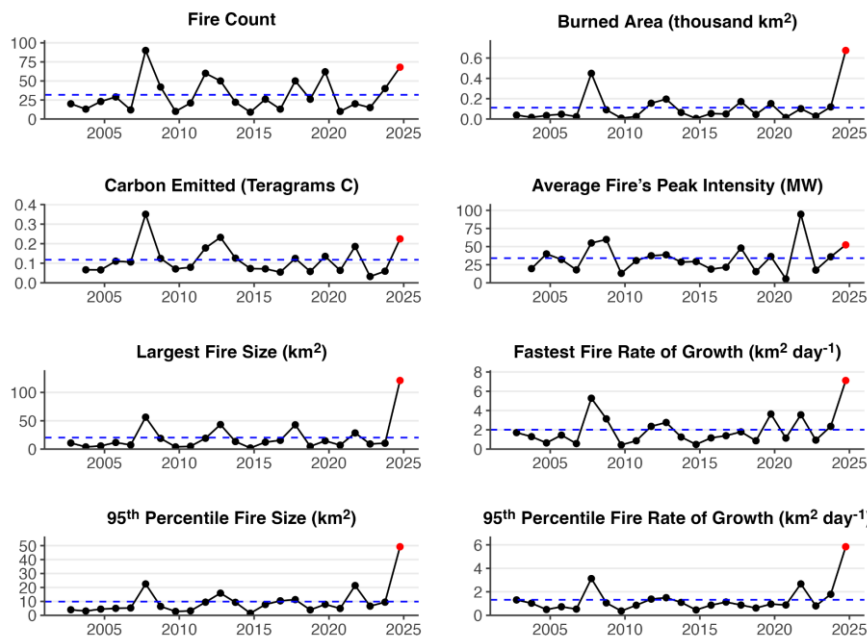


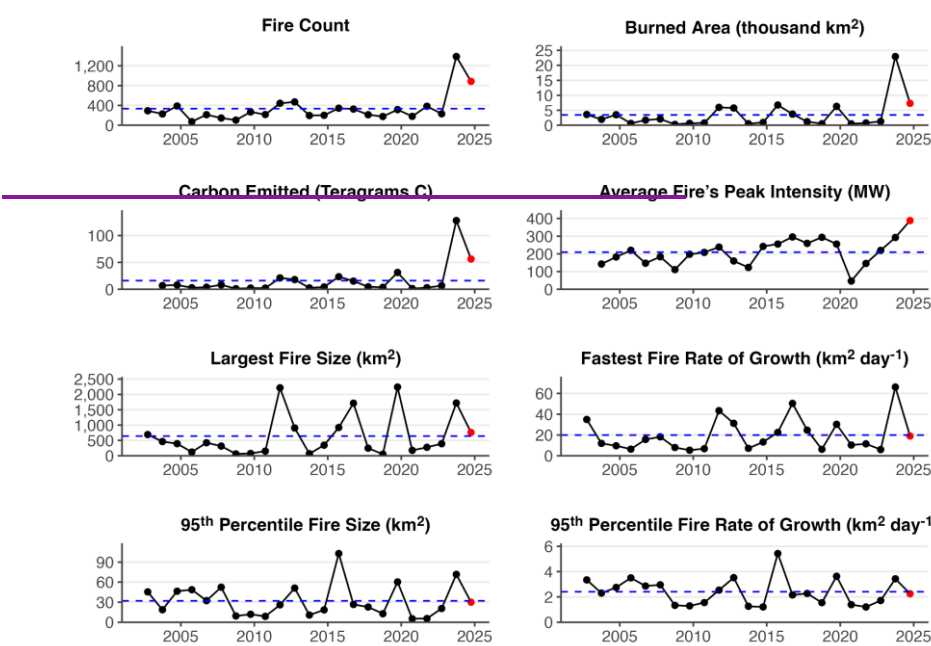
Figure A2: Summary of the 2024-2025 fire season in North Macedonia, as in **Figure A1**.

A4. North America

Wildfires across North America were characterized by above average activity in Canada and the United States and a record-breaking season in Mexico in 2024-2025. This included multiple wildfires that resulted in substantial impacts to human communities, including the Eaton and Palisades wildfires, which are among the most destructive in California, USA history. Following a record breaking 2023 wildfire year, during which almost 150,000 km² burned, Canada once again experienced an above average wildfire season in 2024. A total of 5,686 wildfires burned approximately 46,000 km², marking the six-highest area burned since 1972 based on national records (Skakun et al. 2024). In the United States, over 64,000 wildfires burned over 36,000 km² in 2024, exceeding both the previous 5 and 10-year averages (NICC, 2024). The USA also recorded the second-highest number of Level 4 and 5 National Fire Preparedness days since 1990, reflecting elevated national fire suppression resource commitment associated with high potential for continued emerging wildfires (NICC, 2024). National fire records for Mexico suggest that the country experienced more than 8,000 wildfires in 2024, setting a record for area burned - over 16,500 km² - since record keeping began in 1998 (Comisión Nacional Forestal, 2025), though this record is not reflected in the global datasets compiled as part of this report.

Much of Canada experienced earlier-than-normal snowmelt in 2024, resulting in an early onset of the wildfire season. For example, parts of Alberta experienced snowmelt up to 30 days earlier than average. Drought conditions, which were prevalent across the country in 2023, persisted into 2024 in much of western Canada. Holdover fires from 2023, which smouldered through the winter in northern British Columbia, Alberta (**Figure A3**), and parts of the Northwest Territories, reignited in early spring due to warm and dry conditions (Kolden et al. 2025). This contributed to above average area burned and wildfire emissions in May

4902 (Copernicus Atmosphere Monitoring Service, 2024). Wildfires in May led to evacuations of
4903 Fort Nelson, British Columbia and Fort McMurray, Alberta - an area previously affected by
4904 Canada's costliest wildfire in 2016 (Canadian Forest Service, 2025).



4905

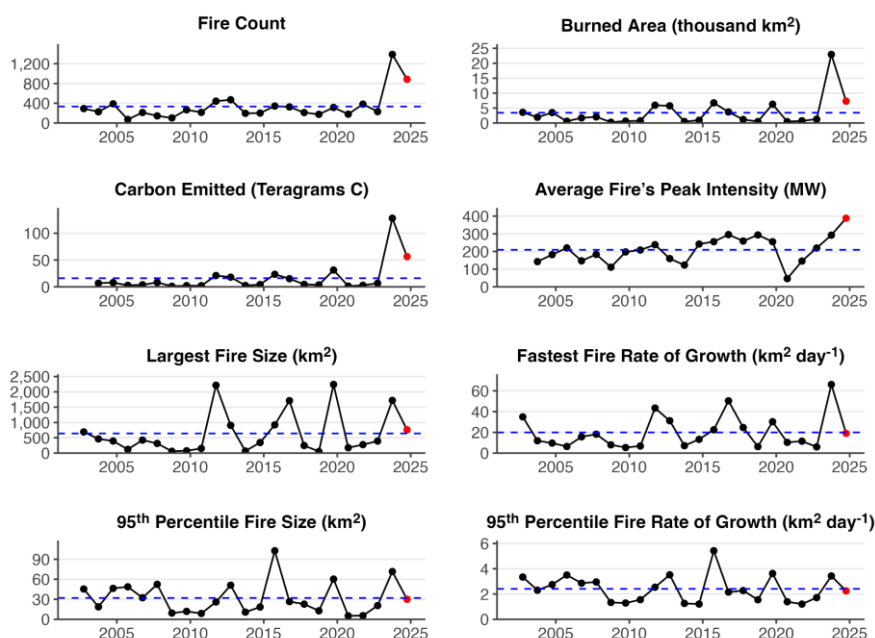


Figure A3: Summary of the 2024-2025 fire season in Alberta, Canada, as in **Figure A1**.

Most of the USA was characterized by normal to high precipitation at the start of 2024, with minimal wildfire activity (NICC, 2024). A heat wave at the end of February 2024 in the southern plains combined with strong winds and high fine fuel loads led to multiple large wildfires, including the record-breaking Smokehouse Creek Wildfire in the Texas Panhandle and western Oklahoma that burned over 4000 km² and resulted in two fatalities before reaching 100% containment in March (Texas House of Representatives Investigative Committee on the Panhandle Wildfires, 2024). Wildfire risk in the southern plains remained elevated for several weeks. Warm and dry conditions in March led to an increase in activity in the central Appalachians region of the eastern US, with the Virginia Department of Forestry reporting over 100 wildfires in 48 hours. By early April, fire activity peaked for the spring fire season in the southern and eastern US. Dry and windy conditions prompted significant growth of large wildfires burning in New Mexico; however, fire activity remained below average in the USA in May (NICC, 2024).

Wildfires in Mexico started increasing in March during Mexico's typical wildfire season. Warm and dry conditions helped to fuel hundreds of wildfires (Comisión Nacional Forestal, 2025; NASA Earth Observatory, 2024b) contributing to Mexico's record breaking wildfire season with anomalously high carbon emissions. Wildfire numbers peaked by mid-March through early May (Comisión Nacional Forestal, 2025).

Wildfire activity remained high across Canada during the summer of 2024, with many regions experiencing above-average area burned. Areas including New Brunswick in the east, and the Northwest Territories, recorded area burned totals among the top five highest since 1972 (Skakun et al., 2024). Hot and dry conditions in July resulted in wildfires forcing the evacuation of Labrador City in Newfoundland and Labrador and John D'Or Prairie and Fox Lake in Alberta

4931 (Canadian Forest Service, 2025). In late July during a period of extreme 99th percentile fire
4932 weather, a fast-moving wildfire resulted in the evacuation of the town of Jasper, Alberta and
4933 destroyed 358 structures resulting in an estimated \$1.23 billion in damages - the second
4934 costliest wildfire in Canadian history (Kolden et al. 2025; Insurance Bureau of Canada, 2025).
4935 There were two fatalities in July related to fire suppression operations in Alberta and the
4936 Yukon.

4937 Large fires continued to burn in northern regions of British Columbia, Alberta, and throughout
4938 the Northwest Territories throughout August and into the fall, resulting in the fourth, sixth, fifth
4939 highest area burned for these areas, respectively, since 1972 (Skakun et al., 2024). Significant
4940 fire activity also developed in Saskatchewan, Manitoba, and Ontario in August, and New
4941 Brunswick experienced the second highest area burned since 1972 (Skakun et al. 2024). In
4942 total, 91 wildfire-related evacuations took place across Canada during the 2024 season,
4943 affecting 56,000 people (Canadian Forest Service, 2025). According to estimates from the
4944 Copernicus Atmosphere Monitoring System, the 2024 wildfire season in Canada produced the
4945 second-highest total emissions recorded since 2003 - surpassed only by the record-breaking
4946 2023 season (Parrington & Di Tomaso, 2024).

4947 Wildfire activity began to pick up in the USA during the later part of June, with multiple fires in
4948 New Mexico resulting in several hundred structures burned (NICC, 2024), two fatalities, and
4949 over \$1 billion in damages (NCEI, 2025). By July, an extreme and long-lasting heatwave
4950 across the western US spurred numerous large wildfires, including the Park Fire in Northern
4951 California that drove thousands to evacuate and destroyed over 700 structures (CALFIRE,
4952 2025). Record breaking dry conditions in Oregon and Washington led to over 100 human-
4953 caused wildfires by early July (US Forest Service, 2024), contributing to a record setting year
4954 in BA and anomalously high carbon emissions in Oregon (**Figure S47**). Through July and
4955 August, hot and dry weather drove numerous large wildfires in the northwestern front range,
4956 including the Stone Canyon wildfire in Colorado that resulted in one fatality and multiple
4957 burned homes and the Remington wildfire in Wyoming that killed hundreds of cattle. During
4958 September, numerous dry lightning strike wildfires occurred in the northwestern US along with
4959 multiple wildfires in southern California associated with extreme heat, including the Airport Fire
4960 that resulted in 22 injuries and 194 damaged structures (CALFIRE, 2025).

4961 Fall was anomalously warm and dry across much of the continental US, with 87% classified
4962 as abnormally dry or in drought by early November (NICC, 2024). The northeast US
4963 experienced hundreds of wildfires that interacted with densely populated regions in October
4964 and November coincident with record-dry conditions and warm temperatures across multiple
4965 states. For instance, New York City experienced its highest number of recorded wildfires
4966 during a two-week period, with every borough experiencing multiple wildfires. The conditions
4967 were unseasonable, with Connecticut, Massachusetts, and Rhode Island setting record red
4968 flag days, despite typical peaks for red flag days occurring in spring (NOAA, 2024).
4969 Massachusetts experienced its most active fall fire season in over 40 years (NICC, 2024). Two
4970 fatalities and hundreds of structures were destroyed before rainfall associated with an
4971 extratropical cyclone halted the fall fire season in the northeast in late November.

4972 Wildfire activity remained low at the end of 2024 and beginning of 2025, except in southern
4973 California. Southern California is climatically prone to experiencing a downslope (katabatic)
4974 wind during the late autumn and winter months known locally as Santa Ana winds. Historically,
4975 the most devastating wildfires in California have occurred when a delayed onset of autumn
4976 precipitation coincides with a Santa Ana wind event (Kolden and Abatzoglou, 2018), but such
4977 concurrences are increasing in frequency with climate change (Goss et al., 2020). In
4978 November and December, Santa Ana wind events produced wildfires that burned nearly 100
4979 km² (10,000 ha) and destroyed over 250 structures, however, this was just a precursor. In
4980 January 2025, the most disastrous wildfire event in modern US history occurred in Los
4981 Angeles County, California. Prolonged drought conditions, unseasonably warm winter

temperatures, and exceptionally powerful Santa Ana winds exceeding 140 km/h created extreme fire weather conditions (Barnes et al., 2025; CNN, 2025). Fire potential was also exacerbated by anomalously wet winters for two years prior, which increased the fine fuel load in the region. The potential for extreme wildfires to develop under dry downslope winds was predicted several days in advance, including by the National Interagency Fire Center (NIFC), the National Weather Service (NWS), and the Storm Prediction Center (SPC; see summary by Wikipedia, 2025) as well as by specialist commentators (e.g. Swain, 2025).

The two most destructive fires-Palisades and Eaton-that burned during the event occurred in the same general locations as destructive fires in 1961 and 1993 during other Santa Ana wind events. These two fires resulted in numerous outcomes with widespread and severe consequences. Among the most devastating were the high fatalities and extensive structure loss. Over 11,500 homes were severely damaged or destroyed across Los Angeles County, and at least 3031 lives were lost, according to the Los Angeles County Coroner (2025). Specifically, the Palisades Fire severely damaged or destroyed nearly 8,000 structures5,614 homes, while the Eaton Fire impacted over 10,000 (CALFIRE, 2025; Wikipedia, 2025). The fires also triggered mass evacuations. At the peak of the crisis, at least 153,000 people were forced to evacuate, with up to 200,000 under evacuation warnings or orders (USGS, 2025b; NPR, 2025; Wikipedia, 2025).

In addition to human displacement and infrastructure damage, the fires severely affected both air and water quality. Air pollution reached hazardous levels, contributing to negative health outcomes for thousands. During the fires, peak PM_{2.5} levels reached 483 µg/m³, an order of magnitude greater than the 35 µg/m³ daily standard set by the US Environmental Protection Agency, resulting in a prolonged period of hazardous air quality (California Air Resources Board, 2025). Over 400 excess deaths in Los Angeles County have since been attributed to exposure to poor air quality during January 2025, including due to lung or heart conditions exacerbated by smoke or stress or to indirect causes (e.g. disruptions to health systems or mental health impacts; Paglino et al., 2025). Municipal water supplies were similarly impacted, with water considered unsafe for tens of thousands of residents in the burned areas for several weeks following the fires and eight water districts in LA county issuing "do-not-use" or "do-not-drink" warnings (City of Pasadena, 2025). Beyond Los Angeles, the political fallout from the crisis led to federal orders to release over 8.3 million cubic meters of water from federal reservoirs further north in California. However, this water did not flow to southern California and was instead vital for irrigating crops in the state's heavily agricultural Central Valley (Levin et al., 2025).

The economic consequences were equally severe. Total economic losses were estimated at US\$140 billion, factoring in property destruction, health costs, business disruption, and infrastructure damage, making this one of the most costly wildfire events in US history (LAEDC, 2025; UCLA Anderson School of Management, 2025). Wider economic disruption is also projected, with estimated losses of US\$4.6-8.9 billion in economic output over five years, 25,000-50,000 job-years lost, and reductions in labour income of US\$1.9-3.7 billion (LAEDC, 2025). The Palisades and Eaton fires directly affected nearly 2,000 businesses (LAEDC, 2025). Moreover, as Los Angeles hosts the largest port on the US Pacific coast, these fires disrupted broader supply chains connected to the Port of LA (ASU, 2025).

Insured losses added another layer of financial strain, with industry estimates ranging from \$20 billion to \$75 billion (PreventionWeb, 2025; Morningstar DBRS, 2025; Insurance Insider, 2025; UCLA Anderson, 2025). This placed substantial pressure on the already volatile home insurance market in California, as well as on most global re-insurers.

The fires also deepened Southern California's ongoing housing and affordability crisis. Thousands of affordable housing units were lost, worsening the existing housing shortage, displacing large numbers of low-income residents, and exacerbating the region's

homelessness problem (Urban Land Institute, 2025; UCLA Anderson, 2025; Vox, 2025). This led to ripple effects, with mass displacement into surrounding communities and beyond in the months that followed (NYT, 2025).

Finally, the aftermath of the fires brought additional physical hazards in the form of debris flows. Given southern California's geology, the region is highly susceptible to erosion and debris flows following wildfires. Several such events occurred after high-intensity rainfall in the weeks following the fires, causing further damage and prompting hundreds of additional evacuations in and near the recently burned areas (USGSa, 2025).

A5. Oceania

Oceania experienced relatively moderate levels of fire during the 2024-25 fire season, although there were still a series of high profile and high impact events across the region. Overall, however, the season did not reach the magnitude of the previous year, which ranked among the top 5 years for BA in Australia since 2002. Where fires occurred and had impacts, lightning and sustained dryness were prominent drivers (Bureau of Meteorology, 2024; Dowdy and Brown, 2025).

The 2024-25 fire season in Western Australia was characterised by record-high temperatures, variable rainfall, and significant soil moisture deficits in coastal areas of the South, Southwest, and West. Over 1,000 large fires burned about ~~470,000~~ ~~ha~~ ~~4,700 km²~~, many in coastal shrubland and woodland over the ~800 km stretch from Gingin, north of Perth, to Carnarvon. The largest fire by area burned occurred near Cervantes in November, where fire ignited by a car crash went on to burn more than ~~80,000~~ ~~ha~~ ~~800 km²~~ and severely impact local honey production. In the inland Goldfields region at Skeleton Rocks, more than ~~44,000~~ ~~ha~~ ~~440 km²~~ of Mallee-heath vegetation of the Great Western Woodlands were burned (according to the Department of Fire Emergency Services (DFES), Rui Feix, pers. comm.). This fire reached extreme intensity, impacting fire-sensitive species and post-fire regeneration cycles in an ecosystem that requires long intervals to recover. A lithium mine in the area was also directly impacted by the fire. Four other large incidents were recorded in the shrublands of the Great Western Woodlands, further affecting these vulnerable ecosystems. Between December and March, numerous fires occurred in the grasslands of the Wheatbelt and Esperance, as well as in the forests of the Perth Hills. These included fires that collectively destroyed seven residential properties in areas east of Mundaring, Arthur River, Wooroloo, and Waroona. In February and March, lightning ignited several large bushfires in native forests and coastal shrubland around Manjimup. Some of these fires burned for up to five weeks and affected more than ~~42,000~~ ~~ha~~ ~~420 km²~~, including areas of Shannon and D'Entrecasteaux National Parks (DFES, Rui Feix, pers. comm.). These incidents required significant aerial support and personnel deployments, including interstate assistance.

Above average rainfall was recorded in Central Australia, leading to expectations of strong grass fuel growth and another period of increased fire activity, after last year's above average season (Verhoeven et al. 2020; Ruscalleda-Alvarez et al. 2023). By the end of October 2024, over ~~5.7 million~~ ~~ha~~ ~~57 thousand km²~~ had burned, much of it stemming from an intense band of dry lightning stretching from the Northern Territory into Queensland in October (according to Northern Territory Fire and Emergency Services, Maggie Towers, pers. comm.). Many of these fires combined with a particularly large fire complex near Devil's Marbles Conservation Reserve (~~450,000~~ ~~ha~~ ~~4,500 km²~~) (NTFES, Maggie Towers, pers. comm.). The fire threatened hotels and other infrastructure and caused temporary closure of a major highway. In late January 2025, a bushfire swept through the West MacDonnell Ranges, affecting approximately ~~80,000~~ ~~ha~~ ~~800 km²~~ across the Tjoritja/West MacDonnell National Park, Standley Chasm, the Tyurretye and Iwupataka Aboriginal Land Trusts, as well as nearby pastoral properties (NTFES, Maggie Towers, pers. comm.). Standley Chasm and sections of the

5081 Larapinta Trail were closed for several days while a 10-day multi-agency effort worked to
5082 contain the fire.

5083 Queensland's north west saw heightened fire activity during Spring, with fire fighters
5084 responding to 40 incidents in Mount Isa alone. One of these fires burned for nearly two months,
5085 reaching over ~~100,000 ha~~ 1 thousand km² according to Queensland Fire Department, (Russell
5086 Stephens-Peacock, pers. comm.). The fires caused an increase in hospital admissions due to
5087 respiratory illnesses and impacted mining operations, pastoral property and Lawn Hill National
5088 Park. The fires affected the habitat and food sources of endangered species such as the
5089 Carpentarian Grass wren, found only in north-western Queensland.

5090 In 2024-25, eastern Australia, comprising southern Queensland, New South Wales (NSW),
5091 and the Australian Capital Territory (ACT), experienced a notably warm period, with significant
5092 rainfall variation across regions and seasons. Although temperatures were above average in
5093 the austral spring, many areas received above average rainfall, thereby reducing fire
5094 occurrence and impacts. Repeated dry lightning started a number of complexes of fires in
5095 remote and difficult to access terrain across NSW, including areas like Lithgow, the
5096 Hawkesbury, Bulga and around Tamworth. Despite the number of fires, NSW saw more
5097 moderate fire weather than other parts of the country and Emergency Warnings were only
5098 issued for three fires (according to New South Wales Rural Fire Service, David Field, pers.
5099 comm.).

5100 The south to south east of Australia (including the states of South Australia, Victoria and
5101 Tasmania) experienced record dryness in the leadup to the fire season. Fires in Chappelvale
5102 and Casterton-Edenhope in late spring signaled an early start to the fire season in Victoria. In
5103 December a band of dry lightning ignited a number of fires including several in the Grampians
5104 National Park. About ~~75,000 ha~~ 750 km² burned in the Grampians, affecting culturally and
5105 ecologically sensitive areas. The coincidence of the fire with Christmas and the peak holiday
5106 season led to major tourism losses and extensive community evacuations. This fire was
5107 contained by January 6 but later in January another band of dry lightning passed through the
5108 west of the state, this time affecting the western side of the Grampians burning another almost
5109 ~~60,000 ha~~ 600 km² (according to Country Fire Authority, Sarah Harris, pers. comm.). By
5110 season's end over two thirds of this important National Park was impacted by fire. Another
5111 significant fire occurred on December 26, a public holiday, in Little Desert National Park in the
5112 state's west. This fire was an extremely fast-moving fire with approximately ~~65,000 ha~~ 650 km²
5113 burning in less than eight hours and a final area burned of ~~90,000 ha~~ 900 km² (according to
5114 Country Fire Authority, Sarah Harris, pers. comm.). These fires required interstate
5115 deployments to assist in the fire fight. The fire season concluded with challenging fires that
5116 burned through rugged terrain in the Gippsland area, impacting the World Heritage-listed
5117 Budjim National Park with its significant cultural heritage. Several planned burns escaped
5118 during the season, highlighting the significant dryness of the area.

5119 In South Australia dry lightning storms in early February combined with severe drought
5120 conditions to cause the Wilmington fire, which burned approximately half of Mount Remarkable
5121 National Park. Firefighting efforts reduced the impact to human, ecological and cultural assets.
5122 Lightning storms in February and March also caused an above average number of fires in
5123 eastern parts of South Australia. While impacts were limited, firefighting resources were
5124 strained. (South Australia, Country Fire Service, Simeon Telfer pers. comm.)

5125 Tasmania faced a significant bushfire season, with up to ~~100,000 ha~~ 1 thousand km² burned
5126 in the state's northwest, including sensitive ecosystems such as the Tarkine rainforest and the
5127 alpine vegetation around Cradle Mountain (according to Tasmania Fire Service, Chris Collins,
5128 pers. comm.). Sparked by intense lightning storms in remote and rugged terrain, the fires
5129 required interstate support to assist with firefighting efforts. The blazes led to evacuations,

threatened heritage sites and caused major disruptions to local businesses and the tourism industry.

In New Zealand the 2024-2025 fire season was moderate, with a couple of minor fires at the end of the 2023/24 fire season (Mar-Jun 2024) and a few more significant fires during the 2024/25 fire season (Jul 2024-Feb 2025). A key feature was the occurrence of a couple of significant wetland fires that burned large areas of peatland (~~2,271~~^{2,434} ~~ha~~^{km²}) and damaged flora and fauna habitat. These fires occurred at Whangamarino Wetland, Waikato (central North Island) in October 2024 and Tiwai Peninsula, Southland (southern South Island) in late January 2025, with both fires just over ~~1,000~~¹⁰ ~~ha~~^{km²}. The fires followed two major peatland fires in 2022 at Kaimaumau in the far north (~~2,434~~²⁴ ~~ha~~^{km²}), and Awarua in the south (~~890~~⁹ ~~ha~~^{km²} and close to this season's Tiwai fire) (according to Fire and Emergency New Zealand, Grant Pearce, pers. comm.). Carbon emissions are likely to be high, given the 2022 fires were estimated to release more than 620,000 t CO₂ (Pronger et al., 2024). There were a number of other noted fires in a mixture of vegetation types including in Canterbury, Northland and North Otago. However, unlike recent years, there were no major house loss incidents, with just a few homes and outbuildings lost across the multiple fires.

A6. South America

The 2024-25 fire season was a remarkable year for fire in South America, with seven of its 13 countries reaching new records in BA since 2002 and widespread records in the fire size, growth rate and intensity distributions (**Figure 3; Figure 4**). Anomalies in BA commenced early in 2024 and persisted through November in some regions (**Figure S4**). As discussed in Section 2.2.2 and Sections 4-6, intense drought and fire weather affected much of South America during the 2024-25 fire season and this drought occurred at a time when socioeconomic factors are increasingly cited as drivers of shifting fire regimes and timing. The event is part of a trend towards an earlier onset of the fire season since 2020, with new record fire counts set for the months of March to May in 2020 and for January in 2022, based on monitoring by Brazil's National Institute for Space Research (INPE) since 1998 (INPE, 2025). During 2024, January, February and June presented the second highest value on record (previous record during 2003 for January and 2007 for the other months, respectively). Fires have expanded into new territories, driven by a combination of climate variability, shifting land-use practices, and governance challenges, as discussed in the study cases, below.

Across South America, the number of fire hotspots recorded by the Queimadas/INPE system (~~511-thousand,000~~⁵²³ ~~thousand,000~~^{thousand,000} hotspots) (INPE, 2025). Compounding climate and human drivers likely led to a widespread extreme fire year across the continent in 2024-2025. The land use fire dependent practices, associated with new deforestation frontiers during an extreme drought year amplified the fire crisis. Amidst rising socioeconomic and environmental impacts of fires in the region, researchers have been calling on governments across the globe to rethink strategies for combating the root causes of extreme wildfires, from climate change to fire-free agricultural practices (UNEP, 2022). Increases in extreme droughts with already vulnerable forest due to extreme climatological events are expected and therefore controlling ignition sources are the only immediate measure for preventing 2024-like scenarios. In this context, major fire events in terms of largest fire size emerged in many parts of Brazil, in Peru, Ecuador and Bolivia during the 2024-25 fire season, with unprecedented levels of BA and exceptional fire weather (**Figure S2**).

In Brazil, one of the most intense droughts in decades, combined with the expansion of the agricultural frontier in Amazonas and Pará states (Santos et al., 2023), caused fires to persist nearly year-round (**Figure S2.4**). In northern Brazil, including much of the Amazon biome, several states such as Amazonas and Pará experienced their largest BA on record. Other states, including Mato Grosso, São Paulo, and Paraná, recorded their highest fire extent in a

single year. In the Pantanal biome, Mato Grosso do Sul experienced the second-largest BA extent on record, the fourth in rank in fire size and fifth regarding the fastest growth. This resulted in estimated losses to agribusiness caused by the fires amounting to R\$ 1.2 billion (~\$222 million) (Câmara, 2024). In addition, Pantanal recorded particulate matter concentrations of 903.2 µg/m³ in September 2024 (Viana et al., 2024), which is 60 times higher than the World Health Organization (WHO) standards. Efforts to contain the flames lasted 78 days and involved the National Force, local communities, environmental organizations, and state fire brigades (Nunes, 2025). The response faced significant challenges, particularly in remote border areas with difficult access and complex logistics. Providing support to isolated populations was especially difficult, with reported cases of respiratory illnesses worsened by smoke exposure, as well as emotional distress, including stress and anxiety (Nunes, 2025).

São Paulo and Mato Grosso state, both centres of large-scale crop production, experienced the fourth-largest BA extent on record. Regarding fire intensity, 2024 was the record for São Paulo, Paraná, Mato Grosso do Sul, Rio de Janeiro and Roraima, and the second in the rank for Amazonas and Goiás. All together, from the southeast to the north of the country, records in one or more fire metrics were observed during this period, placing Brazil in a state of fire emergency.

In general, early fire season onset and long duration occurred across most Brazilian regions, with the first month of anomalous fire ranging from March in most of the Amazonian states and extending through to December. In fact, more fire hotspots were detected in February and March 2024 than in any year since 1998 (INPE, 2025). Record fire counts were observed across states covering more than half of Brazil's territory, and represented a threat almost during the entire year, posing challenges for managing the wildfires response and combat. By August 2024, the National Centre for Early Warning of Natural Disasters (CEMADEN, 2024) pointed out that the drought, covering Amazonia to the southeast, initiated in the second half of 2023, was already one of the strongest in decades. Data from the National Secretariat for Civil Protection and Defence (S2ID, 2024), in December 2024 pointed out that there were 21 of the 27 states with a recognized decree either in state of emergency or calamity due to the drought, affecting more than 520 municipalities in the country. These conditions brought widespread devastation across Brazil in 2024, impacting urban and rural communities and affecting an estimated 18.9 million people nationwide (CNM, 2024). Fire disaster forced 10,700 people from their homes, resulting in housing instability and severe disruptions to livelihoods. Thousands more were affected by the breakdown of essential services, such as school closures (CNM, 2024). Although Brazil does not have an official database on wildfire-related fatalities, existing records point to a rising death toll (Carvalho et al., 2025). Estimates have identified 186 deaths between 2020 and 2024, with 38 in 2024 alone. However, the actual number is likely higher due to underreporting.

Notably, the state of São Paulo in Brazil recorded 8,712 hotspot fires, the highest number since 1998 (INPE, 2025). August and September together accounted for approximately 70 % of these detections (6,134), roughly four times the 1998-2023 August average (914 hotspot fires) and three times the corresponding September average (848 hotspot fires). According to a study by the Amazon Environmental Research Institute (IPAM, 2024), of the 2,600 hotspots fires recorded in the state of São Paulo between August 22 and 24, 81% were in areas of agricultural use - drawing attention to the fact that the state recorded, on the 23rd August alone, more hotspots than the entire Amazon biome. In an even more alarming interval, analysed images from the geostationary satellite indicate that the smoke columns in western São Paulo appeared in just 90 minutes, between 10:30 AM and 12:00 PM on the 23rd August, and, on that same day, the number of fires jumped from 25 to 1,886 hotspot fires, reinforcing the hypothesis of orchestrated action and the unprecedented intensity of these fires.

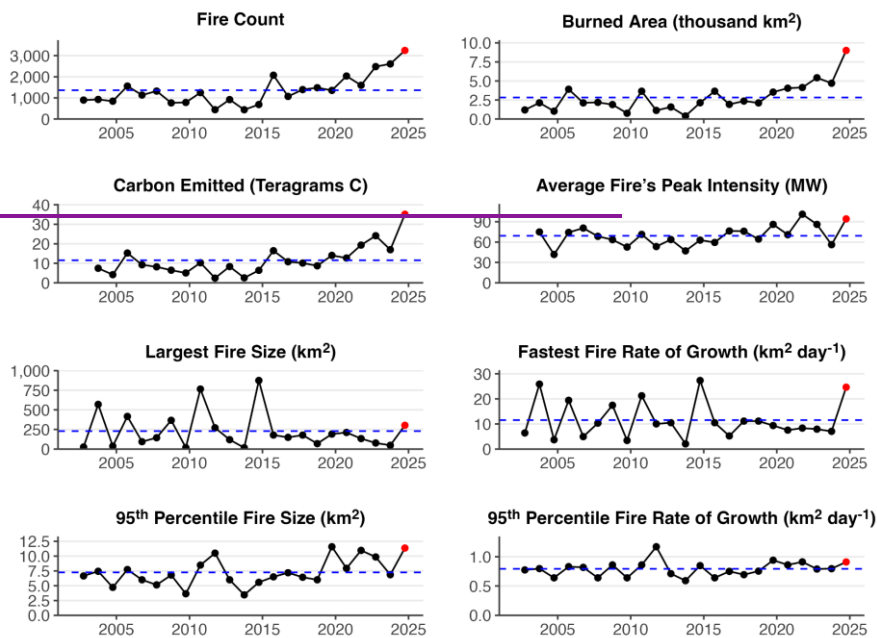
Amazonas state, the largest of Brazil's Amazonian states, can be pointed out as an epicentre of wildfires and its impacts. During 2024, it was ranked as first in number of fires (**Figure 4**) and presented a historical record of fire occurrence in June, July and August, consecutively,

5233 since the monitoring began in 1998 (INPE, 2025). Moreover, it was the 3rd year with the fastest
5234 fire growth rates, with a fire season lasting for 8 months. It has been estimated that fires
5235 affected ~~over 790,000 ha~~ **around 8 thousand km²** of forests, approximately 39% of the affected
5236 area, especially in the southern region of the state (Alencar et al., 2022). The Amazonas state
5237 has been facing an increase in the deforestation rate since 2021, mainly in the southern
5238 region, following the pressure and political speech of Brazil's BR-319 Highway paving. The
5239 lack of governance, associated with illegal logging, land grabbing and public lands invasion,
5240 are some of the drivers of the fire peaks observed in the region (Fearnside, 2022). Moreover,
5241 it has been estimated that the population from this region has been exposed to aerosols
5242 emitted from the wildfires causing a pollution of up to 113 µg/m³, 653% above the 15 µg/m³
5243 standard set by the WHO, according to the data from the Atmosphere Monitoring Service
5244 (CAMS) and Copernicus Climate Change Service (C3S).

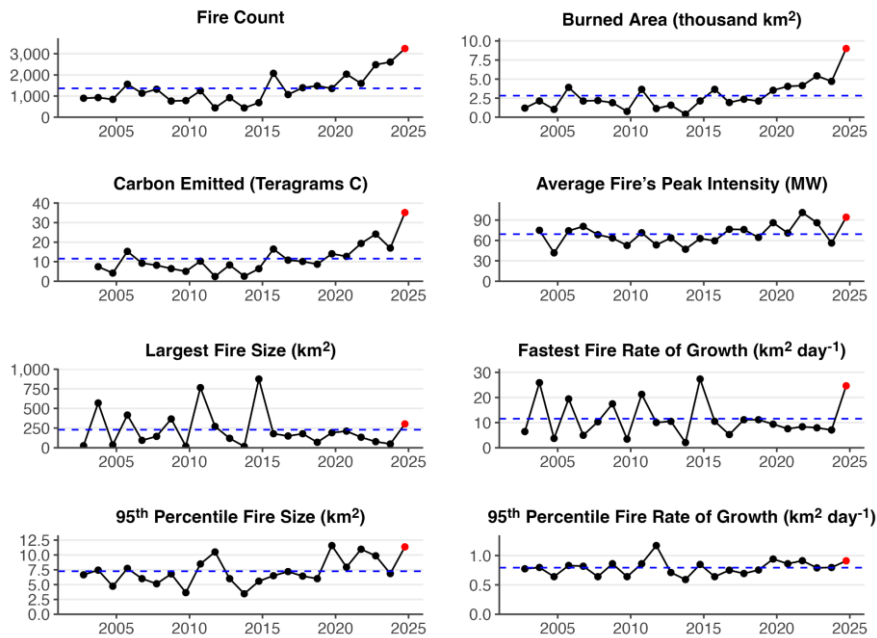
5245 A recent report from the Global Forest Watch (World Resources Institute, 2025) also showed
5246 widespread high levels of forest loss (stand-replacing fire extent) to wildfire in 2024 in the
5247 Amazon biome (including both Brazil and neighbouring Amazonian countries). The highest
5248 rates of forest loss since 2016 was observed, with total forest loss more than doubling in 2024
5249 versus 2023 and 60% of those losses were attributed to wildfires. Note that Global Forest
5250 Watch data define "forest loss" as the complete removal of tree canopy, including areas
5251 affected by stand-replacing fires, but do not capture more subtle or partial fire-related
5252 degradation. As a result, the data may overestimate deforestation while underestimating
5253 degradation, limiting understanding of the broader ecological impacts of wildfires on forests.
5254 Moreover, Indigenous communities were disproportionately affected by wildfires in 2024, a
5255 year that recorded the highest number of fires in territories inhabited by isolated Indigenous
5256 peoples (COIAB, 2024). In 2024, fires in Indigenous lands in Brazil increased by 81%
5257 compared to 2023, accounting for the largest share of Amazonia fires at 24% (Alencar et al.,
5258 2024). In Roraima, uncontrolled fires in indigenous lands have degraded air quality, ravaged
5259 crops, homes, and native vegetation leading to food and water insecurity (ISA, 2024). The
5260 fires have further worsened the ongoing humanitarian crisis in the Yanomami Territory, Brazil's
5261 largest Indigenous land. Local organizations estimate that at least 70,000 people across urban
5262 and rural communities were affected by the lack of access to clean water, a result of the
5263 compounded impacts of drought and fire (WWF-Brazil, 2024).

5264 The implications of extreme fire activity in Amazonia extend beyond immediate ecological
5265 damage. As a globally significant carbon sink and a key part of the terrestrial hydrological
5266 cycle, the Amazon stores an estimated 100-120 Pg of carbon (Malhi et al., 2006). Intensified
5267 fire regimes risk accelerating forest degradation, potentially triggering a biome-scale shift from
5268 net carbon sink to a significant carbon source, releasing several petagrams of carbon and
5269 exacerbating global warming through positive feedbacks (Gatti et al., 2021). Fire-driven
5270 environmental degradation also poses public health risks and economic instability. Biomass
5271 burning increases respiratory illness, especially among populations exposed to prolonged
5272 smoke (Campanharo et al., 2019, 2021). Economically, fire reduces agricultural productivity,
5273 damages infrastructure, and undermines regional development, compounding poverty and
5274 inequality. Costs extend to firefighting programmes and personnel (Morello et al., 2020), as
5275 well as hospitalisations from respiratory or other fire-related conditions (Machado-Silva et al.,
5276 2020). Rising fire activity may also weaken the effectiveness of forest conservation and
5277 restoration policies, including those tied to international climate agreements, threatening long-
5278 term mitigation and adaptation efforts.

5279



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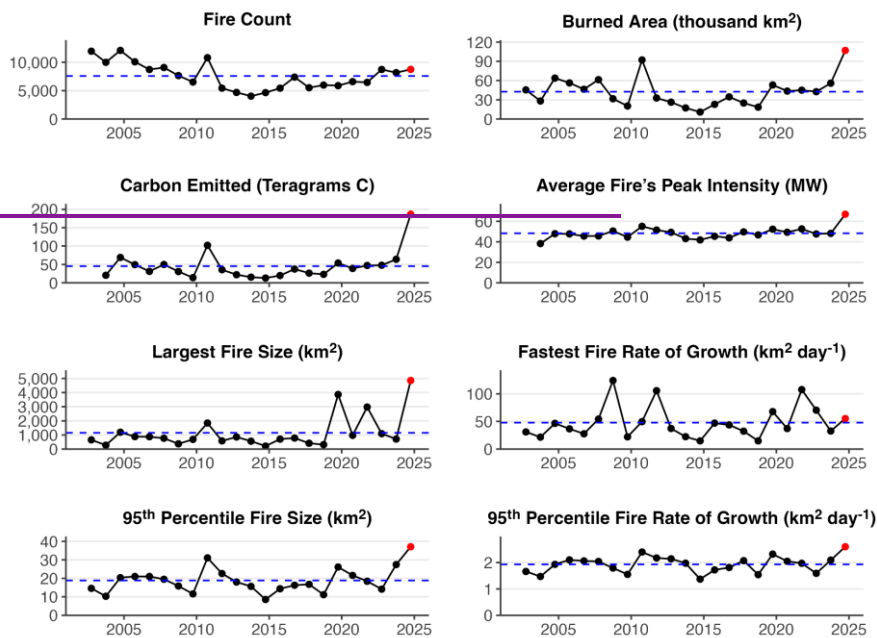
5281 **Figure A4:** Summary of the 2024-2025 fire season in Brazil's Amazonas State, as in **Figure**
5282 **A1.**

5283 Bolivia endured one of its worst fire seasons on record by many measures, intensified by the
5284 El Niño phenomenon, record temperatures, and accelerating deforestation and contributing
5285 significant carbon emissions to the atmosphere (**Figure A5**). These conditions intensified fire
5286 outbreaks, especially in ecologically vulnerable regions such as the Chiquitania and
5287 Amazonian lowlands (Ruiz, 2025). The cumulative number of fire hotspots in 2024 was
5288 923,464, with 77% occurring in Santa Cruz (Chiquitano dry forest), 19% in Beni (Amazonian
5289 lowlands), 1.6% in La Paz, and the rest in other departments including Pando (north
5290 Amazonia) (CEJIS, 2024). A recent report from Global Forest Watch (World Resources
5291 Institute, 2025), found that forest loss in Bolivia tripled in 2024 versus 2023 and was many
5292 times over the annual mean since 2002, with 60% of those losses related to wildfires. The
5293 forest fires have been attributed to a combination of one of the most severe droughts on record
5294 as well as a number of socioeconomic factors and government policies that encourage
5295 agricultural expansion, such as the lifting of soy and beef export quotas, removal of import
5296 taxes on agrochemicals and machinery (World Resources Institute, 2025).

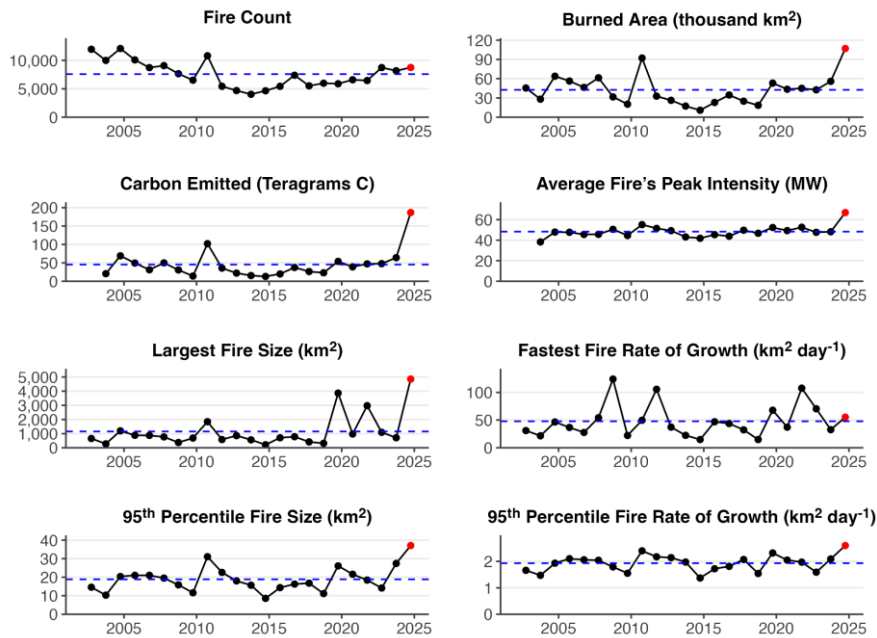
5297 The ~~number of ha~~area burned was a contentious issue in the country. According to the NGO
5298 Fundación Tierra (2024), in September 2024 about ~~10 million ha~~100,000 km² were burned.
5299 The wildfires extended up to November. In early January 2025 an independent group of
5300 experts of the national journal (El Deber, 2025) reported that ~~14 million ha~~140 thousand km²
5301 burned based on the MODIS Terra sensor. In April 2025, the ministry of environment officially
5302 reported ~~12.6 million ha~~126 thousand km² burned in 2024, 12% of the country's territory, with
5303 57% corresponding to primary forest and 43% pastures and agricultural lands (Ministerio de
5304 Medio Ambiente y Agua, 2025). Although wildfires have been occurring regularly in Bolivia
5305 over the past decade, the events of 2024 have been the most catastrophic to date. The event
5306 is considered the second megafire after the one that occurred in 2019. Indigenous lands,
5307 protected areas, and fiscal lands were among the most impacted categories. The Global
5308 Forest Watch cited a lack of early warning systems and adequate firefighting resources as a
5309 factor contributing to high rural exposure to fire and urban exposure to wildfire smoke (World
5310 Resources Institute, 2025). An investigation by Fundación Tierra (2025) reports that wildfires
5311 in Bolivia are mostly intentional, with 66% being maliciously set and 34% resulting from out-
5312 of-control slash-and-burn agricultural practices.

5313 The Bolivian Air Contamination Index reached 537 in the city of Cobija, northern Bolivia (Silva
5314 Trigo, 2024), corresponding to a PM_{2.5} concentration of over 500 µg/m³ (24-hour average), a
5315 level considered extremely harmful and impactful to the health of millions of people in the
5316 region and beyond. As a result, the government declared a sanitary emergency. In addition to
5317 the extensive environmental destruction and incomparable biodiversity loss, these fires have
5318 significantly increased atmospheric carbon emissions, exacerbating regional and global
5319 climate challenges. It is important to note that laws and regulations in Bolivia encourage
5320 agricultural and livestock expansion and are lenient towards the use of fire (Yifan He et al.
5321 2025). Encroachment and illegal land occupation are also pointed to as causes of provoked
5322 wildfires in Bolivia. Efforts in the legislative branch to prohibit or amend these regulations have
5323 not been successful thus far. Therefore, there is a looming risk that similar events may occur
5324 again in the near future.

5325



5326



5327 **Figure A5:** Summary of the 2024-2025 fire season in Bolivia, as in **Figure A1**.

5328 In early 2024, Venezuela experienced its most intense wildfire season on record, with over
5329 30,000 active fires between January and March (NASA FIRMS, 2025). Unlike Brazil,
5330 Venezuela's peak fire season runs from December to April, driven by the northward shift of
5331 the Intertropical Convergence Zone (ITCZ; Katz & Giannini, 2010; Ramírez & Gómez, 2021),
5332 which were intensified by the 2023-2024 El Niño, one of the strongest in decades, creating
5333 an extreme fire weather window (NOAA CPC, 2024). Fires have historically been concentrated
5334 in the Orinoco Llanos, a savanna-dominated region covering approximately one-third of
5335 Venezuela, where fire is used for agricultural purposes and grazing (Bilbao et al., 2020). More
5336 recently, deforestation in tropical forests south of the Orinoco has fueled large fires, like those
5337 seen in 2019 (Lizundia-Loiola et al., 2020). In 2024, wildfires impacted nearly all ecosystems,
5338 from Amazonian humid forests in Bolívar and Amazonas (5,600+ fires, including Canaima
5339 National Park), to flooded savannas in Apure, cloud forests in Henri Pittier, and an estimated
5340 ~~36,000 ha~~ **360 km²** of Caribbean pine lost in Uverito, Latin America's largest plantation (Ciudad
5341 CCS, 2024; Lozada, 2024). Since 2019, Venezuela's National Parks Institute (INPARQUES)
5342 has promoted an intercultural Integrated Fire Management (IFM) strategy, coordinated by an
5343 intersectoral team involving government agencies, local communities, and researchers (Bilbao
5344 et al., 2022). With support from FAO and RAMIF (under ACTO), this national IFM system aims
5345 to strengthen fire management efforts in response to Venezuela's vast ecological and
5346 territorial complexity, as well as to the increasing extreme fire weather conditions projected for
5347 the region, including higher temperatures, prolonged dryness, and lower humidity (Feron et
5348 al., 2024).

5349 A fundamental challenge in the wildfire crisis affecting Bolivia and Venezuela is the complexity
5350 of managing fires in border regions. Many of the most affected areas are located along
5351 international boundaries, where coordination between neighboring countries is often
5352 inadequate or inefficient. The lack of standardized protocols, difficulties in sharing real-time
5353 information, and disparities in firefighting capacities create significant logistical and operational
5354 challenges. Fires in these areas are particularly difficult to control due to overlapping
5355 jurisdictions and administrative barriers that delay response efforts. This is also the case in
5356 other regions in south america, such as the trinational frontier with Acre, Peru and Bolivia
5357 (Pismel et al., 2023) and at the Pantanal region. Without improved cross-border collaboration,
5358 enhanced communication channels, and harmonized fire management strategies, these
5359 transboundary wildfire zones will remain highly vulnerable, exacerbating the broader crisis in
5360 South America.

5361 Ecuador presented the peak in BA during 2024, with an anomaly of 166%, the highest on
5362 record. Official governmental reports from the National secretariat of risk management, there
5363 were almost 6,000 wildfires, ~~83 thousand ha~~ **830 km²** of burned vegetation, 1,663 affected
5364 people, 47 people hurt, 6 deaths, 45,000 animals killed and over 5,000 animals affected,
5365 according to the National Secretariat for Risk Management (SitRep., 2024). These events
5366 were attributed to the extreme drought and land use and land cover conversion fire dependent
5367 practices.

5368 In 2024, there were 13,400 hotspots in Peru, which was 1,000 more than in 2023 (Caceres et
5369 al. 2024). Of the total, 49% were in natural areas such as forests or other natural covers, 35%
5370 in non-forest vegetation types, and 12% in anthropogenic areas. The maximum number of
5371 hotspots occurred in September. According to Caceres et al. (2024), in August and
5372 September, the regions most impacted by fire were Ucayali, Madre de Dios, Huánuco, San
5373 Martín, and Loreto, all belonging to the Amazonia region. In November 2024, the number of
5374 wildfires totaled 1,798, with over ~~80,000 ha~~ **8 thousand km²** burned, 35 people and a countless
5375 number of animals died in the events (Castillo, 2024, Informe Defensorial n.º 225). The
5376 severity of the wildfires was reflected by Castillo (2024) using information from INDECI
5377 (Institute of Civil Defense).

5378 In the extreme south of South America, fires in Patagonia started in early 2025, continuing a
5379 recent trend that aligns with an 80% increase in BA since 2002. In Argentina, the 2024-25 fire
5380 season was the most destructive in decades for northern Patagonia. By late February 2025,
5381 more than ~~30,000~~ [3 thousand km²](#) had burned across Río Negro and Neuquén provinces,
5382 primarily affecting Lanín and Nahuel Huapi National Parks (Greenpeace, 2025). Extreme fire
5383 behavior was driven by prolonged drought, anomalously high temperatures, and intense
5384 westerly winds. Nearly all ignition sources were anthropogenic, amid conditions of critical fuel
5385 dryness (Greenpeace, 2025). The Patagonia 2024-25 fire campaign represents the most
5386 extensive and intense in decades, underscoring the combined influence of climate extremes
5387 and human pressures.

5388 In Chile, fire occurrence and BA were lower during the 2024-25 fire season than in recent
5389 years. The 2024 season reached a BA of ~~73,834~~ [738 km²](#) compared to the ~~429,103~~ [4291](#)
5390 [km²](#) burned in 2023. However, in February 2024, the Valparaíso Region experienced a record-
5391 setting catastrophic fire associated with extreme weather conditions (high temperatures and
5392 strong winds), affecting wildland-urban interface areas with significant material losses and
5393 more than 30 deaths (González et al. 2024). Central and south-central Chile have experienced
5394 an intense and uninterrupted megadrought since 2010, which has increased the size and
5395 severity of wildfires (Garreaud et al 2017; González et al. 2018; Bowman et al. 2019). Priority
5396 steps to advance solving this problem are restoring and managing forest vegetation and
5397 removing highly flammable forest plantations to move towards less fire-prone landscapes.

5398 8. Competing Interests Statement

5399 SV is a member of the editorial board of Earth System Science Data. [NA is an employee of](#)
5400 [BeZero Carbon Ltd., a carbon credit ratings agency for the voluntary carbon market.](#) The
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5421 10. Author Contributions

5422 **Conceptualisation and Project Administration:** DIK, CAB, FdG, MWJ.

5423 **Data Curation, Resources, and Software:** DIK, CAB, FDG, MWJ, MLFB, EBr, JRM, ZL,
5424 ASIB, EBU, AC, EDT, JE, AJH, SL, GM, YQ, FRS, CBS, MAT-V, DvW, JBW, EW, NA, EC,
5425 LG, MP, MT, GRvdW, SV.

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5429 **Formal Analysis:** DIK, CAB, FDG, MWJ, MLFB, EBr, JRM, ZL, KB, IJMF, LF, AJH, SL, GM,
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5432 **Visualisation:** DIK, CAB, FDG, MWJ, MLFB, JRM, ZL, KB, IJMF, LF, SL, AL, FRS, MAT-V,
5433 RV, JBW, NA.
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5435 **Writing - Original Draft Preparation:** DIK, CAB, FDG, MWJ, MLFB, EBr, JRM, ZL, ASIB,
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5440 **Writing - Review & Editing:** All authors.

5441 11. Data Availability

5442 **Section 2:** BA data from NASA's MODIS BA product (MCD64A1) are extended from Giglio et
5443 al. (2018) and are available at Giglio et al. (2021,
5444 <https://lpdaac.usgs.gov/products/mcd64a1v061/>, last access: 6 August 2025). GFED4.1s fire
5445 C emissions data are extended from van der Werf and are available at
5446 <https://globalfiredata.org/> (last access: 6 August 2025). GFAS fire C emissions data are
5447 extended from Kaiser et al. (2012) and are available at
5448 [https://confluence.ecmwf.int/display/CKB/CAMS+global+biomass+burning+emissions+base
5449 d+on+fire+radiative+power+%28GFAS%29%3A+data+documentation](https://confluence.ecmwf.int/display/CKB/CAMS+global+biomass+burning+emissions+base+d+on+fire+radiative+power+%28GFAS%29%3A+data+documentation) (last access: 6 August
5450 2025). Global Fire Atlas are extended from Andela et al. (2019) and are available at Andela
5451 and Jones (2025, <https://doi.org/10.5281/zenodo.11400061>, last access: 6 August 2025).
5452 Regional summaries of the MODIS BA, GFED4.1s, GFAS, and the Global Fire Atlas are
5453 presented here and are available at Jones et al. (2025,
5454 <https://doi.org/10.5281/zenodo.15525674>, last access: 6 August 2025). Regional summaries
5455 of FWI anomalies are available from Turco et al. (2025,
5456 <https://doi.org/10.5281/zenodo.15538595>, last access: 6 August 2025). Studies utilising our
5457 regional summaries should cite both the current article and the primary reference for the
5458 variable(s) of interest: Giglio et al. (2018) for the MODIS MCD64A1 BA product; Andela et al.
5459 (2019) for the Global Fire Atlas; Giglio et al. (2016) for active fire observations of FRP from
5460 MOD14A1 and MYD14A1; Lizundia-Loiola et al. (2022) for the FireCCIS311 BA product; van
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5462 emissions; Vitolo et al. (2020) for FWI from the ECMWF ERA5 reanalysis. **Section 3:** Regional
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5473 driver assessment data for Northeastern Amazonia and Pantanal & Chiquitano available from
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5479 August 2025). **Section 5 (and subsequent sections):** Historical (1960-2013) HadGEM3-A
5480 are available from the Met Office (2025;

5481 <http://catalogue.ceda.ac.uk/uuid/99b29b4bfeae470599fb96243e90cde3>, last access: 6
 5482 August 2025). ConFLAME NRT attribution outputs are available from (Kelley et al. 2025c;
 5483 <https://doi.org/10.5281/zenodo.15641876>, last access: 6 August 2025). FireMIP / ISIMIP
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 5487 et al., 2025b; <https://doi.org/10.5281/zenodo.15807587>, last access: 10 August 2025). Data
 5488 and scripts used to produce Fire Weather Index (FWI) projections at different global warming
 5489 levels are available from Liu & Eden (2025; <https://doi.org/10.5281/zenodo.15790287>, last
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5491 **12. Code Availability**

5492 **Section 3:** Code for regional summaries of population and physical asset exposure have been
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 5506 <https://doi.org/10.5281/zenodo.16779167> (Lampe & Burton 2025), with methods documented
 5507 in Burton, Lampe et al. (2024). The current version of ibicus, used for JULES-ES bias
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 5509 and is described in detail in <https://ibicus.readthedocs.io/en/latest/> (last access: 6 August
 5510 2025). Model code and evaluation for bias-correction of JULES-ES model output can be found
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