

# Energy-conservation datasets of global land surface radiation and heat fluxes from 2000-2020 generated by CoSEB

Junrui Wang<sup>a, b</sup>, Ronglin Tang<sup>a, b, \*</sup>, Meng Liu<sup>c</sup>, Zhao-Liang Li<sup>a, b, c</sup>

<sup>a</sup> State Key Laboratory of Resources and Environment Information System, Institute of Geographic Sciences and Natural Resources Research, Chinese Academy of Sciences, Beijing 100101, China

<sup>b</sup> University of Chinese Academy of Sciences, Beijing 100049, China

<sup>c</sup> State Key Laboratory of Efficient Utilization of Arable Land in China, Institute of Agricultural Resources and Regional Planning, Chinese Academy of Agricultural Sciences, Beijing 100081, China

\* Authors to whom correspondence should be addressed: tangrl@reis.ac.cn

## Abstract

Accurately estimating global land surface radiation [including downward shortwave radiation ( $SW_{IN}$ ), downward longwave radiation ( $LW_{IN}$ ), upward shortwave radiation ( $SW_{OUT}$ ), upward longwave radiation ( $LW_{OUT}$ ) and net radiation (Rn)] and heat fluxes [including latent heat flux (LE), soil heat flux (G) and sensible heat flux (H)] is essential for quantifying the exchange of radiation, heat and water between the land and atmosphere under global climate change. This study presents the first data-driven energy-conservation datasets of global land surface radiation and heat fluxes from 2000 to 2020, generated by our model of Coordinated estimates of land Surface Energy Balance components (CoSEB). The model that integrates GLASS and MODIS remote sensing data, ERA5-Land reanalysis datasets, topographic data, CO<sub>2</sub> concentration data as independent variables and in situ radiation and heat flux observations at 258 eddy covariance sites worldwide as dependent variables within a multivariate random forest technique to effectively learn the physics of energy conservation~~was renewed with a combination of GLASS and MODIS remote sensing data, ERA5-Land reanalysis datasets, topographic data, CO<sub>2</sub> concentration data, and observations at 258 eddy~~

~~covariance sites worldwide from the AmeriFlux, FLUXNET, EuroFlux, OzFlux,~~  
~~ChinaFLUX and TPDC.~~ The developed CoSEB-based datasets are strikingly  
 advantageous in that [1] they are the first RS-based data-driven global datasets that  
 satisfy both surface radiation balance  ~~$(SW_{IN} - SW_{OUT} + LW_{IN} - LW_{OUT} = R_n)$~~  and heat  
 balance  ~~$(LE + H + G = R_n)$~~  among the eight fluxes, as demonstrated by both the  
 radiation imbalance ratio [RIR, defined as  $100\% \times (SW_{IN} - SW_{OUT} + LW_{IN} - LW_{OUT})/R_n$ ]  
 and energy imbalance ratio [EIR, defined as  $100\% \times (R_n - G - LE - H)/R_n$ ] of 0, [2] the  
 radiation and heat fluxes are characterized by high accuracies, where (1) the RMSEs  
 $(R^2)$  for daily estimates of  $SW_{IN}$ ,  $SW_{OUT}$ ,  $LW_{IN}$ ,  $LW_{OUT}$ ,  $R_n$ ,  $LE$ ,  $H$  and  $G$  from the  
 CoSEB-based datasets at 44 independent test sites were ~~28.51~~37.52  $W/m^2$  ~~(0.81),~~  
~~10.39~~4.20  $W/m^2$  ~~(0.42),~~ ~~14.29~~22.47  $W/m^2$  ~~(0.90),~~ ~~10.62~~3.78  $W/m^2$  ~~(0.95),~~ ~~22.40~~9.66  
 $W/m^2$  ~~(0.77),~~ ~~24.38~~30.87  $W/m^2$  ~~(0.60),~~ ~~22.67~~9.75  $W/m^2$  ~~(0.44)~~ and ~~6.77~~5.69  $W/m^2$   
~~(0.44), respectively, as well as for 8-day estimates were 12.81  $W/m^2$ , 7.08  $W/m^2$ , 9.22~~  
 ~~$W/m^2$ , 8.34  $W/m^2$ , 13.38  $W/m^2$ , 19.99  $W/m^2$ , 17.44  $W/m^2$  and 4.25  $W/m^2$ , respectively,~~  
 (2) the CoSEB-based datasets, in comparison to the mainstream products/datasets (i.e.  
 GLASS, BESS-Rad, BESSV2.0, FLUXCOM, MOD16A2, PML\_V2 and ETMonitor)  
 that generally separately estimated subsets of the eight flux components, better agreed  
 with the in situ observations. Our developed datasets hold significant potential for  
 application across diverse fields such as agriculture, forestry, hydrology, meteorology,  
 ecology, and environmental science, which can facilitate comprehensive studies on the  
 variability, impacts, responses, adaptation strategies, and mitigation measures of global  
 and regional land surface radiation and heat fluxes under the influences of climate  
 change and human activities. The CoSEB-based datasets are open access and available  
 through the National Tibetan Plateau Data Center (TPDC) at  
<https://doi.org/10.11888/Terre.tpdc.302559> (Tang et al., 2025a) and through the Science  
 Data Bank (ScienceDB) at <https://doi.org/10.57760/sciencedb.27228> (Tang et al.,  
 2025b).

**Key words:** Surface energy balance; Surface radiation balance; Shortwave/Longwave

radiation; Net radiation; Sensible/Latent heat flux; Evapotranspiration; CoSEB

## 1 Introduction

Land surface radiation balance and heat balance play important roles in Earth's climate system, representing the physical processes by which the surface-atmosphere absorbs and redistributes radiation and heat fluxes (Berbery et al., 1999; Betts et al., 1996; Mueller et al., 2009; Sellers et al., 1997; Xu et al., 2022a), and facilitating the exchange of water, energy, carbon, and other agents essential to climatic and ecological systems and human society (Jia et al., 2013; Wang et al., 2012; Wild, 2009; Wild et al., 2012; Xia et al., 2006). Accurately monitoring the spatial and temporal variations of global land surface radiation [including downward shortwave radiation ( $SW_{IN}$ ), downward longwave radiation ( $LW_{IN}$ ), upward shortwave radiation ( $SW_{OUT}$ ), upward longwave radiation ( $LW_{OUT}$ ) and net radiation ( $R_n$ )] and heat fluxes [including latent heat flux ( $LE$ ), soil heat flux ( $G$ ) and sensible heat flux ( $H$ )] is indispensable for quantifying the exchange of radiation, heat and water between the land and atmosphere under global climate change (Ersi et al., 2024; Liang et al., 2019; Rios & Ramamurthy, 2022; Tang et al., 2024a; Wang et al., 2021), and for studying solar energy utilization (Tang et al., 2024b; Zhang et al., 2017), hydrological cycle (Huang et al., 2015; Wild & Liepert, 2010), ecosystem productivity (Nemani et al., 2003), agricultural management (De Wit et al., 2005) and ecological protection (Tang et al., 2023). Remote sensing (RS) technology, with its high spatial-temporal resolution and applicability over large areas, is considered to be the most effective and economical means for obtaining global land surface radiation and heat fluxes (Liu et al., 2016; Van Der Tol, 2012; Zhang et al., 2010).

In past decades, numerous RS-based products/datasets of global surface radiation and heat fluxes have significantly advanced, which were generally generated by physical (Li et al., 2023; Mu et al., 2011; Yu et al., 2022) or statistical methods (Jiao et al., 2023; Jung et al., 2019; Peng et al., 2020). However, two key limitations still exist in these products. Firstly, most available products provide only a single component of

land surface radiation or heat fluxes, e.g. ETMonitor (Zheng et al., 2022) and MOD16A2 (Mu et al., 2011) only estimating LE, leading to the failure to satisfy surface radiation balance and heat balance when the single radiation or heat flux is utilized in conjunction with products containing other radiation and heat components (Wang et al., 2025), and further posing significant uncertainties to understand the interactions and redistributions of surface radiation and energy in the Earth-atmosphere system. Secondly, a few products, e.g., FLUXCOM (Jung et al., 2019) and GLASS (Jiang et al., 2015; Zhang et al., 2014), generated datasets for multiple components of surface radiation and heat fluxes by using ~~independent—separate~~ estimates from the uncoordinated models, which make them difficult to abide by surface radiation and heat conservation. These energy-imbalanced and radiation-imbalanced estimates among multiple components from previous products/datasets severely limit their in-depth applications in analyzing the spatial and temporal trends, simulating the physical processes of radiation, heat and water cycles as well as revealing the attributions and mechanisms in Earth-surface system under global climate change. It was ~~impending and~~ imperative to develop global datasets of land surface radiation and heat fluxes characterized by high accuracies, radiation balance as well as heat balance, to better meet the requirements in practical applications of various fields.

Our proposed data-driven model/framework of Coordinated estimates of land Surface Energy Balance components (CoSEB) (Wang et al., 2025), which effectively learns the underlying physical interrelations (i.e., surface energy conservation law) among multiple targeted variables, provides an unprecedented opportunity to develop global datasets of land surface radiation and heat fluxes that can not only simultaneously provide high-accuracy estimates of these components but also adhere to surface radiation- and heat-conservation laws.

The objectives of this study are twofold: (1) to develop high-accuracy datasets of global land surface radiation and heat fluxes, which comply with the principles of radiation balance and heat balance, using our CoSEB model renewed based on in situ

observations, remote sensing data and reanalysis datasets; (2) to validate the datasets/model estimates against data from in situ observations, mainstream products as well as estimates from uncoordinated random forest (RF) techniques. Section 2 introduces the data resources used in this study. Section 3 briefly describes the method we used to estimate global surface radiation and heat fluxes. Section 4 presents the evaluation of the datasets/model estimates generated by our renewed CoSEB model. Section 5 discusses the superiority, potential applications and uncertainties of the developed datasets. Data availability is given in Section 6, and a summary and conclusion is provided in Section 7.

## 2 Data

### 2.1 Ground-based observations

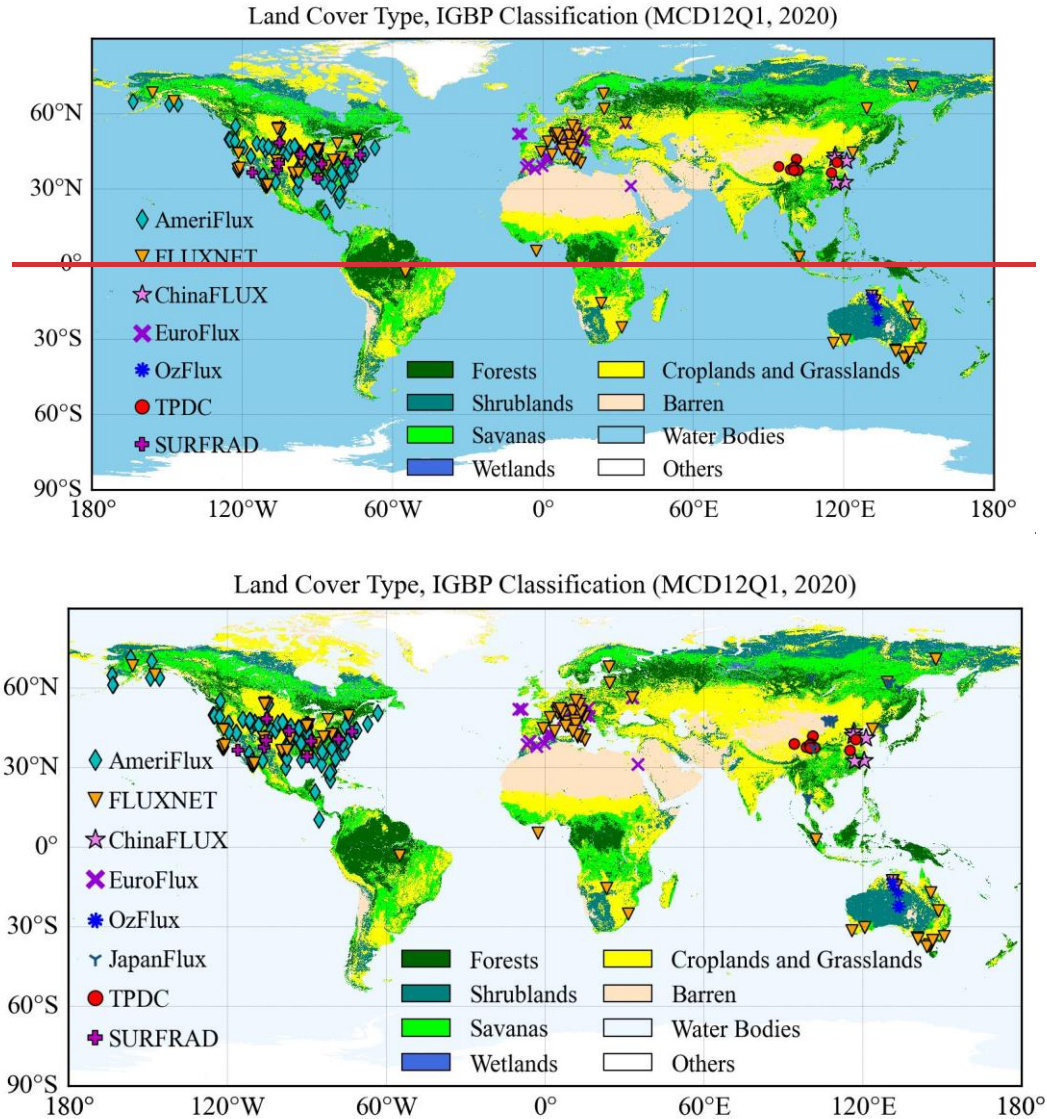
In this study, the in situ observations of land surface radiation and heat fluxes at ~~258–302~~ eddy covariance (EC) sites from the networks of AmeriFlux (~~145–174~~ sites, 2000–2020, <https://AmeriFlux.lbl.gov/Data/>, last access: 6 August 2024), EuroFlux (72 sites, 2000–2020, <http://www.europe-fluxdata.eu/>, last access: 6 August 2024), OzFlux (5 sites, 2007–2012, <https://data.ozflux.org.au/>, last access: 6 August 2024), FLUXNET (108 sites, 2000–2014, <https://FLUXNET.org/Data/download-Data/>, last access: 6 August 2024), JapanFlux (15 sites, 2001–2020, <https://ads.nipr.ac.jp/japan-flux2024/>, last access: 10 October 2025), ChinaFLUX (5 sites, 2005–2020, <http://www.chinaflux.org/>, last access: 6 August 2024) and National Tibetan Plateau/Third Pole Environment Data Center (TPDC, 13 sites, 2012–2020, <https://Data.tpdc.ac.cn/en/Data>, last access: 6 August 2024) were used (Fig. 1), where 37, 48 and 5 sites in FLUXNET were also shared in AmeriFlux, EuroFlux and OzFlux, respectively. These ~~258–302~~ sites were filtered out from all collected ~~1008–1098~~ sites by following the quality-assurance and quality-control steps, including: (1) any site with a missing component of any of the  $SW_{IN}$ ,  $SW_{OUT}$ ,  $LW_{IN}$ ,  $LW_{OUT}$ , LE, H and G was excluded, reducing the ~~1008–1098~~ sites to ~~388–472~~ sites for further analysis; (2) any

half-hour period with missing data for any of these components was excluded; (3) the half-hourly ground-based observations with quality-control flag of 2 or 3 (bad quality) were removed but quality-control flag of 0 and 1 (good quality) were maintained; (4) a daily average of the half-hour observations was calculated for each day with greater than 80% good-quality data, further reducing the ~~388-472~~ sites to ~~286-355~~ sites; (5) the aggregated daily LE and H were corrected for energy imbalance using the Bowen ratio method when the daily energy balance closure [defined as  $(LE + H) / (R_n - G)$ ] varied between 0.2 and 1.8 following Wang et al. (2025) to exclude physically implausible measurements; ~~(56) extreme outliers in the daily evaporative fraction were further removed by excluding values outside the 1st–99th percentile range, a common practice in flux and remote sensing studies~~ (Bartkowiak et al., 2024; Wang et al., 2023), further reducing the 355 sites to 337 sites. outliers were discarded, corresponding to the 1 and 99 quantiles of the daily evaporation fraction, further reducing the 286 sites to 268 sites. Besides, the RS ~~data-products/datasets~~ involved in this study collocated at the sites should not be missing, finally reducing the ~~268-337~~ sites to ~~258-302~~ sites for analysis. Note that the  $R_n$  at these sites used in this study was calculated from the sum of net longwave radiation ( $LW_{IN}$  minus  $LW_{OUT}$ ) and net shortwave radiation ( $SW_{IN}$  minus  $SW_{OUT}$ ), rather than using the observed  $R_n$  directly, to ensure surface radiation balance in training datasets.

These ~~258-302~~ sites used in this study cover a wide range of global climate regimes across 14 land cover types, including (1) evergreen needleleaf forests (ENF, ~~54-55~~ sites); (2) evergreen broadleaf forests (EBF, ~~41-12~~ sites); (3) deciduous needleleaf forests (DNF, ~~1-7~~ sites); (4) deciduous broadleaf forests (~~DBF~~, ~~39-40~~ sites); (5) mixed forests (MF, 8 sites); (6) closed shrublands (CSH, 5 sites); (7) open shrublands (OSH, ~~9-11~~ sites); (8) woody savannas (WSA, 6 sites); (9) savannas (SAV, 10 sites); (10) grasslands (GRA, ~~54-62~~ sites); (11) permanent wetlands (WET, ~~16-22~~ sites); (12) croplands (CRO, ~~43-59~~ sites); (13) water bodies (WAT, 1 sites); (14) cropland/natural vegetation mosaics (CVM, ~~1-4~~ sites). Among them, 44 sites (~15% of the total, see Table S1) were isolated



to serve as spatially independent sites to test the generated datasets and they did not participate in the development of the model/datasets.



**Fig. 1** Spatial distribution of the [258-302](#) eddy covariance sites from AmeriFlux, FLUXNET, EuroFlux, OzFlux, [JapanFlux](#), ChinaFLUX and TPDC, and nine radiation sites from SURFRAD involved for analysis in this study.

Furthermore, ground-based radiation observations from nine sites that are located in large flat agricultural areas covered by crops and grasses from SURFRAD were also introduced to validate land surface radiation estimates. Similar to the preprocessing performed on the observations of the [258-302](#) EC sites, the  $SW_{IN}$ ,  $SW_{OUT}$ ,  $LW_{IN}$ ,  $LW_{OUT}$  and  $R_n$  from the SURFRAD were also quality-controlled and aggregated to daily data.

Spatial distribution of the 258-302 EC sites and nine radiation sites from SURFRAD are shown in Fig. 1, with site details (latitude, longitude, land cover types, digital elevation model and temporal coverage) provided in Supplementary Tables S1 and S2.

**Table 1 Summary of mainstream datasets/products for inter-comparison used in this study**

Products/ datasets	Reso- lution	Time- coverage	Variables	Algorithms	References
GLASS	0.05°/ daily	2000- 2018	$SW_{IN}$ , $LW_{IN}$ , $LW_{OUT}$ , $R_n$	Machine- learning, direct- estimation- algorithm	Wang et al. (2015); Xu et al. (2022b); Jiang et al. (2015)
BESS-Rad	0.05°/ daily	2000- 2020	$SW_{IN}$	BESS process- model	Ryu et al. (2018)
BESSV2.0	0.05°/ daily	2000- 2020	$R_n$ , LE	BESS process- model	Li et al. (2023)
FLUXCOM	0.0833°/ 8-day	2000- 2020	$R_n$ , LE, H	Model tree- ensembles	Jung et al. (2019)
MOD16A2	500-m/ 8-day	2000- 2020	$R_n$ , LE	Modified Penman- Monteith equation Penman-Monteith- Leuning model,	Mu et al. (2011)
PML_V2	500-m/ 8-day	2002- 2020	LE	Priestly Taylor- equation and Gash- model Shuttleworth- Wallace two-	Zhang et al. (2019)
ETMonitor	1-km/ daily	2000- 2020	LE	source-scheme, Gash model and Penman equation	Zheng et al. (2022)

## 2.2 Climate/meteorology and remote sensing data

To generate global datasets of land surface radiation and heat fluxes from 2000 to 2020, five types of climate/meteorology and remote sensing data were used in this study, including:

- (1) ERA5-Land reanalysis datasets (<https://cds.climate.copernicus.eu/>, last access: 6 August 2024) with the spatial resolution of ~9 km from 1950 (Muñoz-Sabater et al., 2021). Following our previous work (Wang et al., 2025), this study used



variables from the ERA5-Land datasets to drive the model, including near-surface 2 m air temperature ( $T_a$ ), soil temperature in layer 1 (0-7 cm,  $T_{s1}$ ), soil volumetric moisture content in layer 1 (0-7 cm,  ~~$SMI$~~ ), solar radiation reaching the surface of the earth ( $SW_{IN}^{ERA5}$ ), net thermal radiation at the surface ( $LW_{net}$ ), pressure of the atmosphere ( ~~$PA$~~ ), 10 m wind speed ( ~~$WS$~~ ), precipitation ( ~~$P_r$~~ ) and the 2 m dewpoint temperature, daily minimum and maximum air temperature [for calculating relative air humidity ( ~~$RH$~~ )].

(2) GLASS datasets (<https://glass.bnu.edu.cn/>, last access: 6 August 2024), which provide the 500 m 8-day leaf area index ( $LAI$ ) and fractional vegetation cover ( $FVC$ ) from February 2000 to December 2021.

(3) MOD44B product (<https://lpdaac.usgs.gov/>, last access: 6 August 2024), which offers yearly 250 m percent tree cover ( $PTC$ ) since 2000, representing the percentage (0~100%) of a pixel covered by tree canopy.

(4) NOAA/GML atmospheric carbon dioxide ( $CO_2$ ) concentration data, providing monthly global marine surface mean data since 1958 ([ftp://aftp.cmdl.noaa.gov/products/trends/co2/co2\\_mm\\_gl.txt](ftp://aftp.cmdl.noaa.gov/products/trends/co2/co2_mm_gl.txt), last access: 6 August 2024).

(5) GMTED2010 topographic data ([https://topotools.cr.usgs.gov/gmted\\_viewer/gmted2010\\_global\\_grids.php](https://topotools.cr.usgs.gov/gmted_viewer/gmted2010_global_grids.php), last access: 6 August 2024), providing 500 m digital elevation model (DEM), slope, and aspect.

The ~9 km ERA5-Land datasets were spatially interpolated to 500 m using the cubic convolution method, and the 250 m  $PTC$  was resampled to 500 m using the arithmetic averaging method.

### 2.3 Mainstream datasets/products for inter-comparison

Mainstream RS-based datasets/products of moderate-resolution global land surface radiation and heat fluxes were collected for inter-comparison (Table 1),

including (1) the daily  $0.05^\circ$  GLASS  $SW_{IN}$ ,  $LW_{IN}$ ,  $LW_{OUT}$  and Rn products from 2000 to 2018 (<https://glass.bnu.edu.cn/>, last access: 6 August 2024), (2) the daily  $0.05^\circ$  Breathing Earth System Simulator Radiation (BESS-Rad)  $SW_{IN}$  products from 2000 to 2020 (<https://www.environment.snu.ac.kr/bess-rad>), (3) the daily  $0.05^\circ$  BESS Version2.0 (BESSV2.0) Rn and LE products from 2000 to 2020 (<https://www.environment.snu.ac.kr/bessv2>), (4) the 8-day  $0.0833^\circ$  FLUXCOM Rn, LE and H products from 2001 to 2020 (<https://fluxcom.org/>, last access: 6 August 2024), (5) the daily 1 km ETMonitor LE product from 2000 to 2020 (<https://data.casearth.cn/>, last access: 6 August 2024), (6) the 8-day 500 m Penman-Monteith-Leuning Version2 (PML\_V2, <https://www.tpdac.ac.cn/>, last access: 6 August 2024) LE product from 2000 to 2020; and (7) the 8-day 500 m MOD16A2 (<https://lpdaac.usgs.gov/>, last access: 6 August 2024) LE product from 2000 to 2020.

The GLASS  $SW_{IN}$  products are derived from a combination of the GLASS broadband albedo product and the surface shortwave net radiation estimates, where the surface shortwave net radiation is estimated using linear regression with MODIS top-of-atmosphere (TOA) spectral reflectance (Wang et al., 2015). The GLASS  $LW_{IN}$  and  $LW_{OUT}$  products are generated using densely connected convolutional neural networks, incorporating Advanced Very High-Resolution Radiometer (AVHRR) TOA reflectance and ERA5 near-surface meteorological data (Xu et al., 2022b). The GLASS Rn products are estimated from the meteorological variables from MERRA2 and surface variables from GLASS using the multivariate adaptive regression splines model (Jiang et al., 2015). The BESS-Rad and BESSV2.0 estimate  $SW_{IN}$  and Rn using a radiative transfer model (i.e., Forest Light Environmental Simulator, FLiES) with an artificial neural network based on MODIS and MERRA2 reanalysis datasets, and using FLiES based on MODIS products and NCEP/NCAR reanalysis data, respectively (Li et al., 2023; Ryu et al., 2018). Moreover, the BESSV2.0 (Li et al., 2023), MOD16A2 (Mu et al., 2011), PML\_V2 (Zhang et al., 2019) and ETMonitor (Zheng et al., 2022) generated global LE by physical models, such as Penman-Monteith equation, Priestley-Taylor

equation and/or Shuttleworth-Wallace two-source scheme. The FLUXCOM Rn, LE and H datasets are obtained through multiple machine learning methods based on in situ observations from FLUXNET and remote sensing and meteorological data (Jung et al., 2019). For better consistency, RF-based 8-day 0.0833° Rn and Bowen ratio-corrected LE and H for the periods of 2000 to 2020 from the FLUXCOM were used in this study.

**Table 1 Summary of mainstream datasets/products for inter-comparison used in this study**

<u>Products/ datasets</u>	<u>Reso- lution</u>	<u>Time coverage</u>	<u>Variables</u>	<u>Algorithms</u>	<u>References</u>
<u>GLASS</u>	<u>0.05°/ daily</u>	<u>2000- 2018</u>	<u><math>SW_{IN}</math>, <math>LW_{IN}</math>, <math>LW_{OUT}</math>, Rn</u>	<u>Machine learning, direct estimation algorithm</u>	<u>Wang et al. (2015); Xu et al. (2022b); Jiang et al. (2015)</u>
<u>BESS-Rad</u>	<u>0.05°/ daily</u>	<u>2000- 2020</u>	<u><math>SW_{IN}</math></u>	<u>BESS process model</u>	<u>Ryu et al. (2018)</u>
<u>BESSV2.0</u>	<u>0.05°/ daily</u>	<u>2000- 2020</u>	<u>Rn, LE</u>	<u>BESS process model</u>	<u>Li et al. (2023)</u>
<u>FLUXCOM</u>	<u>0.0833°/ 8-day</u>	<u>2000- 2020</u>	<u>Rn, LE, H</u>	<u>Model tree ensembles</u>	<u>Jung et al. (2019)</u>
<u>MOD16A2</u>	<u>500 m/ 8-day</u>	<u>2000- 2020</u>	<u>LE</u>	<u>Modified Penman- Monteith equation Penman Monteith- Leuning model,</u>	<u>Mu et al. (2011)</u>
<u>PML_V2</u>	<u>500 m/ 8-day</u>	<u>2002- 2020</u>	<u>LE</u>	<u>Priestly Taylor equation and Gash model</u>	<u>Zhang et al. (2019)</u>
<u>ETMonitor</u>	<u>1 km/ daily</u>	<u>2000- 2020</u>	<u>LE</u>	<u>Shuttleworth- Wallace two- source scheme, Gash model and Penman equation</u>	<u>Zheng et al. (2022)</u>

### 3 Methods

The method used to generate global datasets of land surface radiation and heat fluxes is based on the CoSEB model/framework, which was developed by our recently

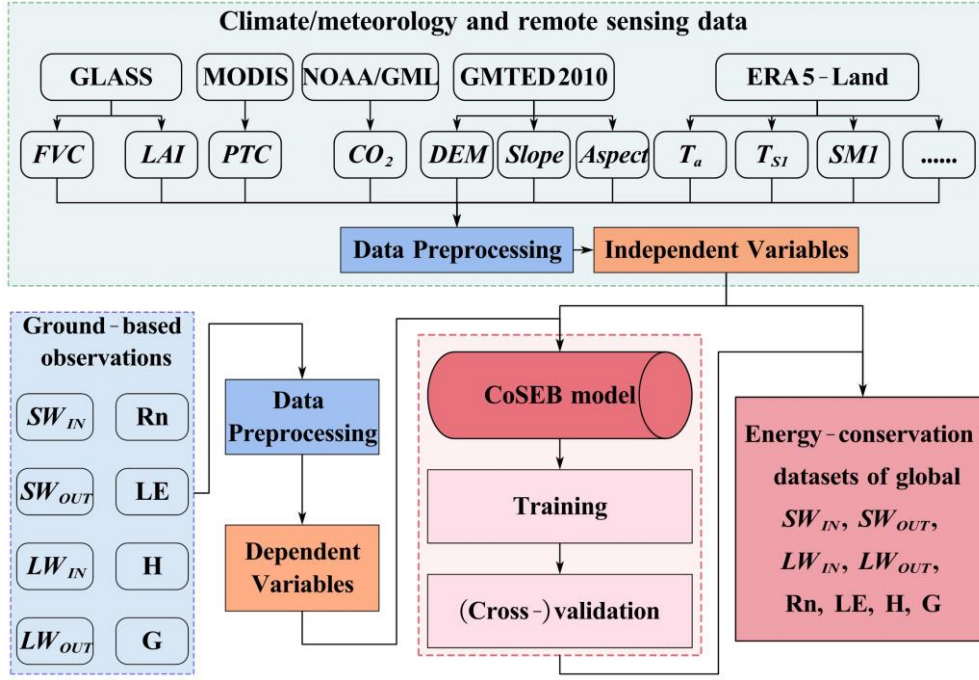
previously published work (Wang et al., 2025), to coordinately estimate global land surface energy balance components (including  $R_n$ ,  $LE$ ,  $H$  and  $G$ ) using the multivariate random forest technique, with a combination of MODIS and GLASS products, ERA5-Land reanalysis datasets, and in situ observations at 336 EC sites ~~from the FLUXNET, AmeriFlux, ChinaFLUX, EuroFlux, OzFlux and Heihe River Basin flux network~~. The CoSEB model was demonstrated to be able to produce high-accuracy estimates of land surface energy components, with the RMSE of  $<17 \text{ W/m}^2$  and  $R^2$  of  $> 0.83$  for estimating 4-day  $R_n$ ,  $LE$  and  $H$ , and the RMSE of  $<5 \text{ W/m}^2$  and  $R^2$  of  $0.55$  for estimating 4-day  $G$ . The most praiseworthy superiority of the CoSEB model lies in its ability to balance the land surface energy components, with an energy imbalance ratio [EIR, defined as  $100\% \times (R_n - G - LE - H) / R_n$   ~~$100\% \times (R_n - G - LE - H) / R_n$~~ ] of 0.

To coordinately estimate land surface radiation and heat fluxes that comply with both radiation balance and heat balance, one of the key procedures in the construction of the CoSEB model was to prepare training datasets that satisfy surface radiation and heat balance. For this purpose, the energy-imbalance corrections on daily in situ observed  $LE$  and  $H$  were conducted by the most widely applied Bowen ratio method [ $H^{corr} = \frac{H}{H + LE} \times (R_n - G)$ ,  $LE^{corr} = \frac{LE}{H + LE} \times (R_n - G)$ , where  $H^{corr}$  and  $LE^{corr}$  represent the sensible heat flux and latent heat flux after energy-imbalance correction, respectively] with the aid of  $R_n$  and  $G$  observations, and the in situ  $R_n$  was calculated from the sum of in situ observed net longwave radiation ( $LW_{IN}$  minus  $LW_{OUT}$ ) and net shortwave radiation ( $SW_{IN}$  minus  $SW_{OUT}$ ). The input variables to renew the CoSEB model include: (1) climate/meteorology:  $T_a$ ,  $SW_{IN}^{ERA5}$ ,  $LW_{net}$ ,  $WS$ ,  $PA$ ,  $P_r$ ,  $RH$ ,  $CO_2$  concentration; (2) vegetation and soil:  $LAI$ ,  $FVC$ ,  $PTC$ ,  $T_{st}$   ~~$T_{st}$~~ ,  $SMI$   ~~$SMI$~~ ; (3) topography data:  $DEM$ ,  $Slope$  and  $Aspect$ , in addition to longitude ( $Lon$ ), latitude ( $Lat$ ), and inverse relative distance from the Earth to the Sun ( $dr$ ), in which the  $dr$  was calculated as  $dr = 1 + 0.033 \times \cos\left(\frac{2\pi \times DOY}{365}\right)$ , where  $DOY$  represents the day of year.

Considering that the footprint of the site-based measurements of turbulent heat fluxes is generally at a scale of hundreds of meters, to reduce the effect of differences of spatial scales between ground-based measurements (dependent variables) and remotely sensed/reanalysis datasets (independent variables), we renewed the CoSEB model at a spatial scale of 500 m for coordinately estimating global daily land surface radiation and heat fluxes, which can be expressed as follows:

$$\begin{pmatrix} SW_{IN}, SW_{OUT}, LW_{IN}, \\ LW_{OUT}, Rn, LE, H, G \end{pmatrix} = f \left( \begin{pmatrix} Lon, Lat, T_a, T_{s1}, SM1, SW_{IN}^{ERAS}, LW_{net}, PA, WS, P_r, dr \\ RH, LAI, FVC, PTC, DEM, Slope, Aspect, CO_2 \end{pmatrix} \right) \quad (1)$$

To enhance model generalization, the renewed CoSEB model was reoptimized using random and grid search methods, resulting in different hyperparameters of decision trees, a maximum depth of 21, and minimum samples split and leaf of 8 from those of Wang et al. (2025). Site-based 10-fold cross-validation was employed to evaluate the transferability and generalization of the CoSEB model by randomly dividing all sites into ten folds, where the samples from each fold of sites in turn served as validation datasets while the remaining folds were used as training datasets, ensuring that the validation was conducted on sites spatially independent from the training data. For comparison, eight RF-based uncoordinated models for separate estimates of  $SW_{IN}$ ,  $SW_{OUT}$ ,  $LW_{IN}$ ,  $LW_{OUT}$ ,  $Rn$ ,  $LE$ ,  $H$  and  $G$  were also constructed using the same inputs as those in the renewed CoSEB model. Site-based 10-fold cross-validation was employed to assess the transferability and generalization of the CoSEB model by randomly dividing all sites into ten folds, where each fold in turn serves as validation datasets while the other folds as the training datasets, ensuring the validation of the estimates of the CoSEB was conducted at sites that are spatially independent from those selected for the training datasets. Furthermore, to benchmark the coordinated estimates from the renewed CoSEB model, eight RF-based uncoordinated models were constructed, each separately estimating one of  $SW_{IN}$ ,  $SW_{OUT}$ ,  $LW_{IN}$ ,  $LW_{OUT}$ ,  $Rn$ ,  $LE$ ,  $H$  or  $G$  using the same inputs as those in the renewed CoSEB model. Fig. 2 illustrates the flowchart for generating global datasets of land surface radiation and heat fluxes by the CoSEB model.



**Fig. 2 Flowchart for generating energy-conservation datasets of global land surface radiation [including downward shortwave radiation ( $SW_{IN}$ ), downward longwave radiation ( $LW_{IN}$ ), upward shortwave radiation ( $SW_{OUT}$ ), upward longwave radiation ( $LW_{OUT}$ ) and net radiation (Rn)] and heat fluxes [including latent heat flux (LE), soil heat flux (G) and sensible heat flux (H)] by the CoSEB model renewed from in situ observations at 258 sites worldwide and collocated remote sensing and reanalysis datasets.**

## 4 Results

### 4.1 Validation of the CoSEB model

#### 4.1.1 Site-based 10-fold cross-validations at 258 EC sites

Fig. 3 and Fig. 4 present the scatter density plots of the site-based 10-fold cross-validation of daily  $SW_{IN}$ ,  $LW_{IN}$ ,  $SW_{OUT}$ ,  $LW_{OUT}$ , Rn, LE, H and G estimated from the renewed CoSEB model and the RF-based uncoordinated models, respectively, by using the validation datasets collected at 258 EC sites worldwide. Results indicated that the estimates from both the CoSEB model and the RF-based uncoordinated models agreed well with the in situ observations, with the coefficient of determination ( $R^2$ ) varying between 0.80 and 0.95 for  $SW_{IN}$ ,  $LW_{IN}$ ,  $LW_{OUT}$  and Rn, and between 0.59 and 0.67 for  $SW_{OUT}$ , LE and H. The CoSEB model, with the root mean square error (RMSE) of 26.82 to 34.25 W/m<sup>2</sup> and mean absolute error (MAE) of 18.83 to 24.49 W/m<sup>2</sup> for  $SW_{IN}$ , Rn, LE and H, the RMSE of 12.24 to 17.75 W/m<sup>2</sup> and the MAE of 8.39 to 13.70 W/m<sup>2</sup> for

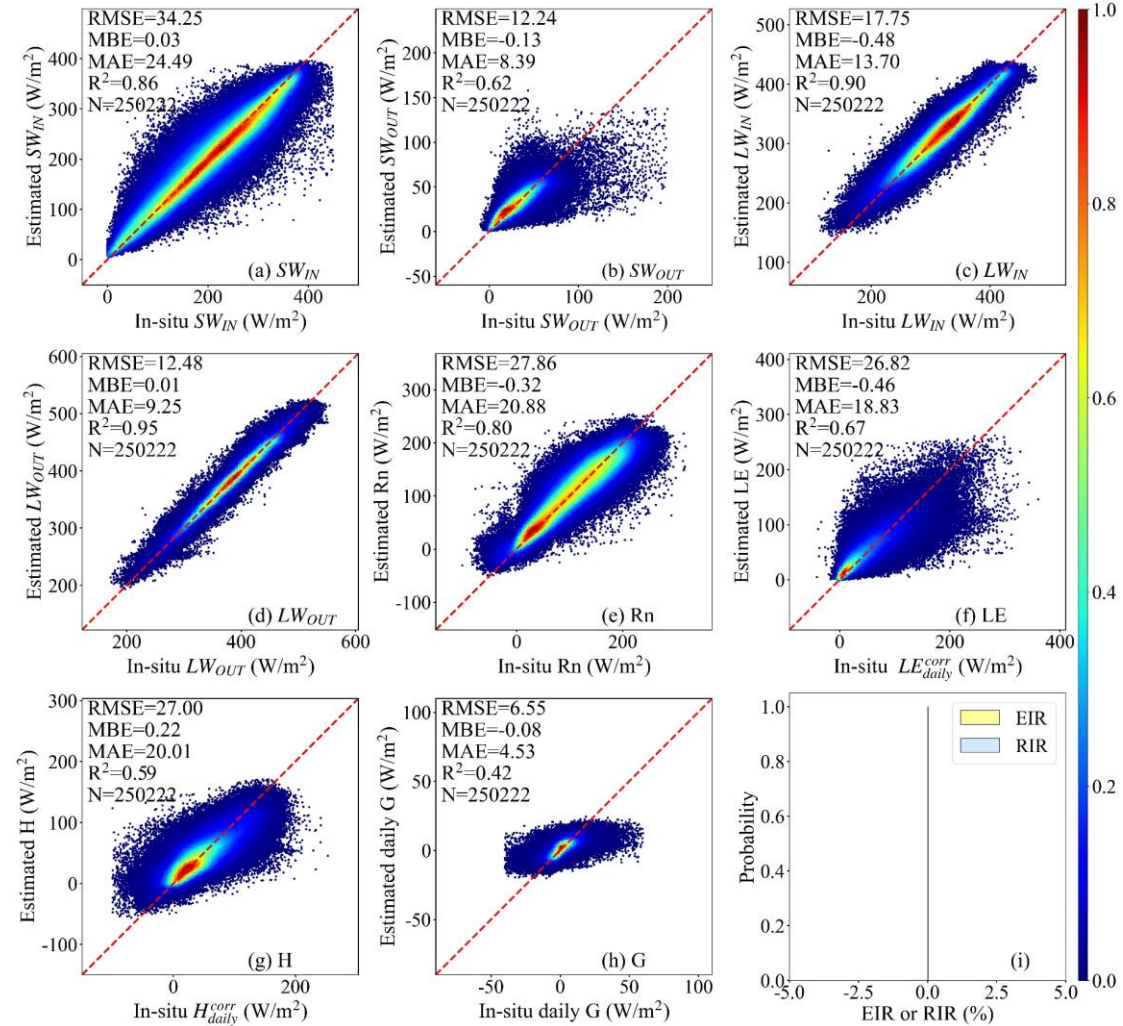


$SW_{OUT}$ ,  $LW_{IN}$  and  $LW_{OUT}$ , demonstrated comparable accuracies to the RF-based models, with the RMSE of 27.07 to 33.34 W/m<sup>2</sup> and MAE of 19.29 to 23.64 W/m<sup>2</sup> for  $SW_{IN}$ ,  $R_n$ ,  $LE$  and  $H$ , the RMSE of 12.12 to 16.93 W/m<sup>2</sup> and the MAE of 8.68 to 12.99 W/m<sup>2</sup> for  $SW_{OUT}$ ,  $LW_{IN}$  and  $LW_{OUT}$ . In the validation of daily  $G$ , both the CoSEB and RF-based models yielded RMSEs below 7 W/m<sup>2</sup>. Comparisons with the corresponding training results (Table S3 in the Supplementary Material) indicated that although the CoSEB model performed better on the training datasets, its overall performance remained stable, suggesting that the CoSEB model was not affected by overfitting.

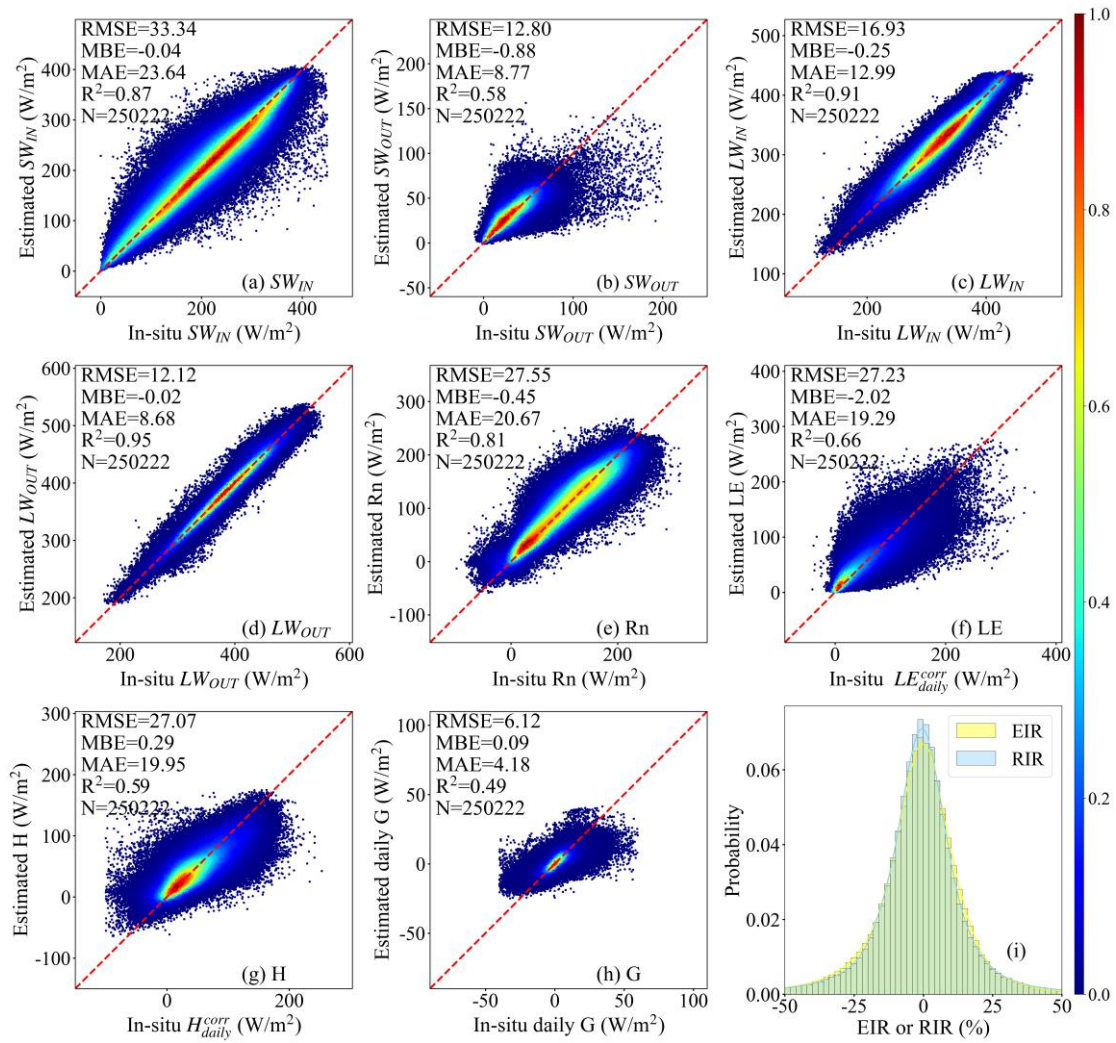
Strikingly, the CoSEB model exhibited large superiority in balancing the surface radiation and heat fluxes, with the radiation imbalance ratio [RIR, defined as  $100\% \times (SW_{IN} - SW_{OUT} + LW_{IN} - LW_{OUT})/R_n$ ], and energy imbalance ratio [EIR, defined as  $100\% \times (R_n - G - LE - H)/R_n$ ] of 0, while the RF-based uncoordinated models showed substantial imbalances of the surface radiation and heat fluxes, with RIR and EIR that were approximately normally distributed, having absolute mean values of 38.84% and 31.22%, respectively, and reaching as high as 50% in some cases. Furthermore, the RIR as well as EIR tended to be higher under lower solar radiation, air temperature, or FVC, with more frequent low values of these three variables leading to a broader and less peaked distribution of RIR and EIR (see Fig. S1 in the Supplementary Material).

~~It should be pointed out that the performances of both the renewed CoSEB model and the RF-based models could be further improved if the site-based 10-fold cross-validation was replaced with the sample-based 10-fold cross-validation (Figs. S1 and S2 in the Supplementary Material). Specifically, for the CoSEB model, using the sample-based 10-fold cross-validation decreased the RMSE by 0.61 to 3.92 W/m<sup>2</sup> for five radiation components and  $G$ , and by 6.25 W/m<sup>2</sup> and 5.50 W/m<sup>2</sup> for  $LE$  and  $H$ , respectively, in comparison to using the site-based 10-fold cross-validation. Likewise, for the RF-based models, the RMSE decreased by 1.41 to 5.25 W/m<sup>2</sup> for five radiation~~

components and  $G$ , and by  $9.63 \text{ W/m}^2$  and  $7.43 \text{ W/m}^2$  for  $LE$  and  $H$ , respectively. The  $R^2$  of both the CoSEB model and the RF-based models using the sample-based 10-fold cross-validation increased by 0.02 to 0.28 compared to the  $R^2$  using the site-based 10-fold cross-validation.



**Fig. 3** Scatter density plots of the site-based 10-fold cross-validation of daily downward shortwave and longwave radiation ( $SW_{IN}$  and  $LW_{IN}$ ), upward shortwave and longwave radiation ( $SW_{OUT}$  and  $LW_{OUT}$ ), net radiation ( $Rn$ ), soil heat flux ( $G$ ), latent heat flux ( $LE$ ) and sensible heat flux ( $H$ ) derived by the CoSEB model against in situ observed  $SW_{IN}$ ,  $LW_{IN}$ ,  $SW_{OUT}$ ,  $LW_{OUT}$ ,  $Rn$ ,  $G$ , and energy imbalance-corrected  $LE$  ( $LE_{daily}^{corr}$ ) and  $H$  ( $H_{daily}^{corr}$ ). The EIR and RIR in the subfigure (i) represent the energy imbalance ratio and radiation imbalance ratio, which are defined as  $100\% \times (Rn - G - LE - H)/Rn$  and  $100\% \times (SW_{IN} + LW_{IN} - SW_{OUT} - LW_{OUT} - Rn)/Rn$ , respectively. The colorbar represents the normalized density of data points.



**Fig. 4 Same as Fig. 3, but for estimates from RF-based uncoordinated models.**

#### 4.1.2 Validation at nine radiation sites from SURFRAD

To further illustrate the generality and transferability of the renewed CoSEB model, the validation of estimates of the five radiation components (including  $SW_{IN}$ ,  $SW_{OUT}$ ,  $LW_{IN}$ ,  $LW_{OUT}$ , Rn) derived from both the CoSEB model and RF-based uncoordinated models against observations at nine radiation sites from SURFRAD was performed, as shown in Fig. 5. The results showed that both the CoSEB model and the RF-based models achieved high accuracy in estimating daily  $SW_{IN}$ ,  $SW_{OUT}$ ,  $LW_{IN}$ ,  $LW_{OUT}$  and Rn, with the RMSE of  $\sim 30 W/m^2$  for  $SW_{IN}$ ,  $\sim 14 W/m^2$  for  $SW_{OUT}$  and  $LW_{IN}$ ,  $\sim 12 W/m^2$  for  $LW_{OUT}$  and  $\sim 24 W/m^2$  for Rn, with the  $R^2 > 0.9$  for  $SW_{IN}$ ,  $LW_{IN}$  and  $LW_{OUT}$ ,  $\sim 0.65$  for  $SW_{OUT}$  and  $\sim 0.85$  for Rn. Compared to the results of the site-based 10-fold cross-validation at 258 EC sites, the performances at nine radiation sites showed slight

improvements, with the RMSE decreasing by 0.74 to 4.54 W/m<sup>2</sup> for  $SW_{IN}$ ,  $LW_{IN}$ ,  $LW_{OUT}$  and  $R_n$  in the CoSEB model, but a slight degradation with the RMSE increasing by ~1.05 W/m<sup>2</sup> for  $SW_{OUT}$ , suggesting the robust performance of the CoSEB model. Furthermore, the CoSEB model demonstrated a large superiority in maintaining surface radiation balance among the five radiation components, with the RIR of 0, in contrast to the RF-based models, which failed to meet this balance, exhibiting significant RIR exceeding 50%.

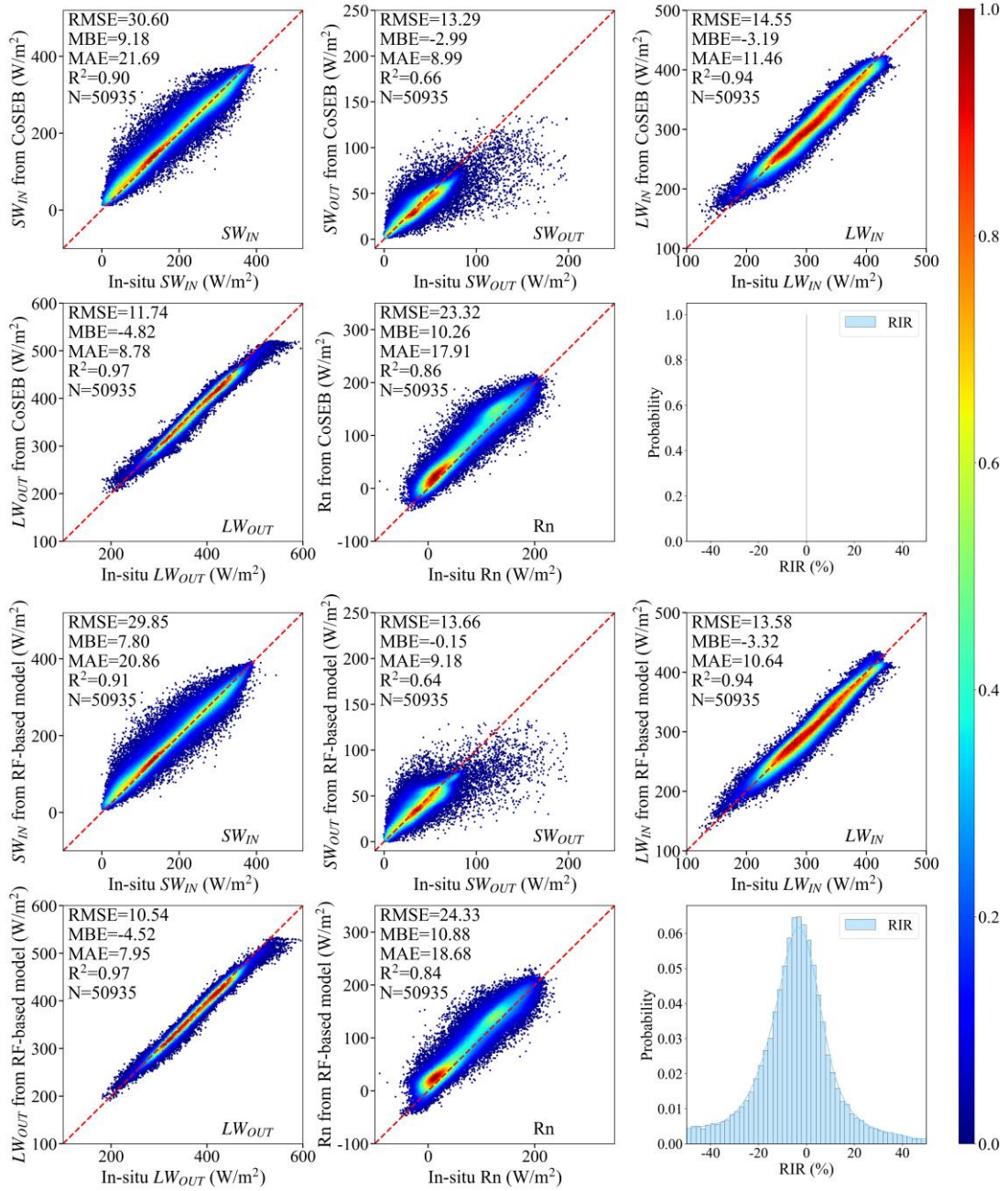
## 4.2 Validation and inter-comparisons of the CoSEB-based datasets

As demonstrated in Section 4.1, the renewed CoSEB model with a spatial scale of 500 m achieved comparable accuracies to the RF-based uncoordinated models but outperformed them in balancing surface radiation and heat fluxes. Evidenced by the validation for its superiority, the renewed CoSEB model was then applied to the spatially aggregated input datasets to generate our developed global daily datasets with a spatial resolution of 0.05°. To further assess the performance of the developed CoSEB-based datasets, in situ observations from another 44 spatially independent test sites (see Section 2.1), which were not involved in model construction and datasets generation, were used for validation. Mainstream products (i.e. GLASS, BESS-Rad, BESSV2.0, FLUXCOM, PML\_V2, MOD16A2 and ETMonitor) were also involved for inter-comparison at the 44 test sites.

Note that due to the lack of moderate-resolution global RS-based products/datasets of daily and/or 8-day  $SW_{OUT}$ ,  $H$  and  $G$ , the intercomparison between different products/datasets was impossible. Instead, we conducted a validation of these components from the CoSEB-based datasets against in situ observations at 44 test sites, as shown in Figs S2 and S3 in the Supplementary Material. Results indicated that the CoSEB-based datasets could provide good estimates of  $SW_{OUT}$ ,  $H$  and  $G$ , with the RMSEs ( $R^2$ ) of 14.20 W/m<sup>2</sup> (0.42), 29.75 W/m<sup>2</sup> (0.44) and 5.69 W/m<sup>2</sup> (0.44) at daily scale, respectively, and the RMSE ( $R^2$ ) of 12.19 W/m<sup>2</sup> (0.39) and 4.60 W/m<sup>2</sup> (0.47) for 8-day  $SW_{OUT}$  and  $G$ , respectively.







**Fig. 5** Scatter density plots of the validation of daily downward shortwave and longwave radiation ( $SW_{IN}$  and  $LW_{IN}$ ), upward shortwave and longwave radiation ( $SW_{OUT}$  and  $LW_{OUT}$ ) and net radiation ( $Rn$ ) from the renewed CoSEB model (upper two rows) and RF-based uncoordinated models (lower two rows) -based-datasets against in situ observations at nine radiation sites from SURFRAD. The RIR represents the radiation imbalance ratio, defined as  $100\% \times (SW_{IN} - SW_{OUT} + LW_{IN} - LW_{OUT}) / Rn$ .

The colorbar represents the normalized density of data points.

#### 4.2 Validation and inter-comparisons of the CoSEB-based datasets

As demonstrated in Section 4.1, the renewed CoSEB model with a spatial scale of

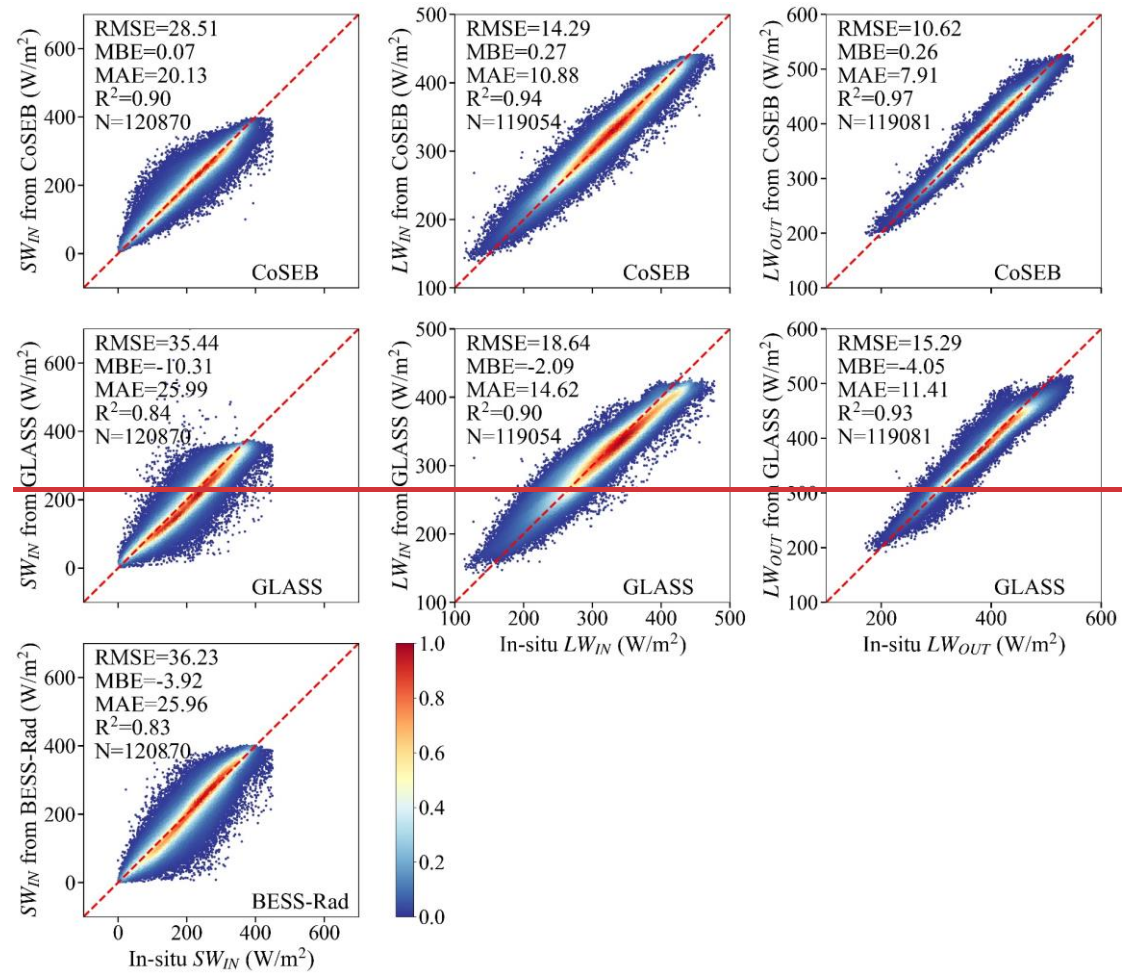


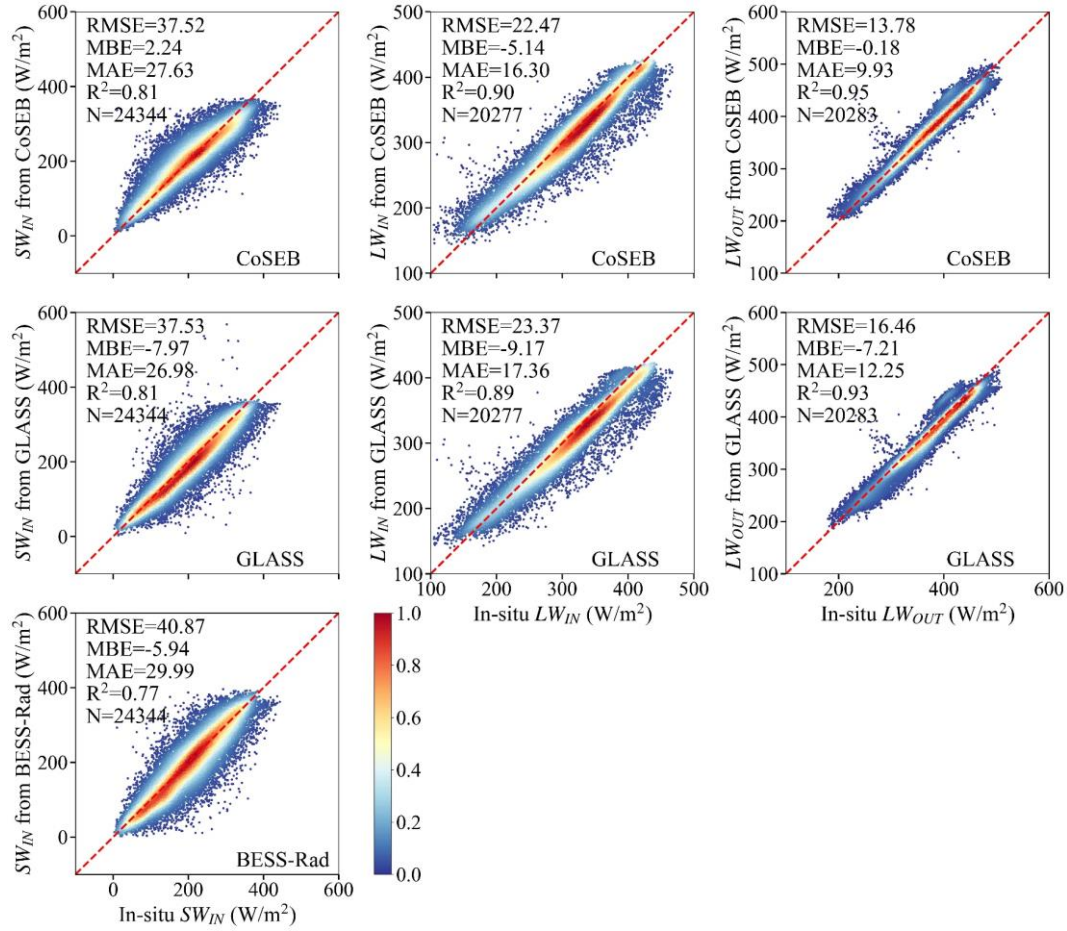
500 m achieved comparable accuracies to the RF-based uncoordinated models but outperformed them in balancing surface radiation and heat fluxes. Evidenced by the validation for its superiority, the renewed CoSEB model was then applied to the spatially aggregated input datasets to generate our developed global daily datasets with a spatial resolution of  $0.05^\circ$ . To further assess the performance of the developed datasets, in situ observations at 134 sites out of the 258 EC sites were further used to test the performance of the CoSEB-based datasets, where the 134 sites were selected based on the commonly applied criterion (Salazar-Martínez et al., 2022; Tang et al., 2024a) that the fraction of the dominant land cover types (from the 500 m MCD12Q1 product) exceeded 80% within the  $0.05^\circ$  grid, ensuring surface homogeneity and spatial representativeness of the observations. Mainstream products (i.e. GLASS, BESS-Rad, BESSV2.0, FLUXCOM, PML\_V2, MOD16A2 and ETMonitor) were also involved for inter-comparison at the 134 EC sites.

Note that due to the lack of moderate-resolution global RS-based products/datasets of daily and/or 8-day  $SW_{OUT}$ ,  $H$  and  $G$ , the intercomparison between different products/datasets was impossible. Instead, we conducted a validation of these components from the CoSEB-based datasets against in situ observations at 134 EC sites, as shown in Figs S3 and S4 in the Supplementary Material. Results indicated that the CoSEB-based datasets could provide good estimates of  $SW_{OUT}$ ,  $H$  and  $G$ , with the RMSE of  $10.39 \text{ W/m}^2$ ,  $22.67 \text{ W/m}^2$  and  $6.77 \text{ W/m}^2$  at daily scale, respectively, and the RMSE of  $7.08 \text{ W/m}^2$  and  $4.25 \text{ W/m}^2$  for 8-day  $SW_{OUT}$  and  $G$ , respectively.

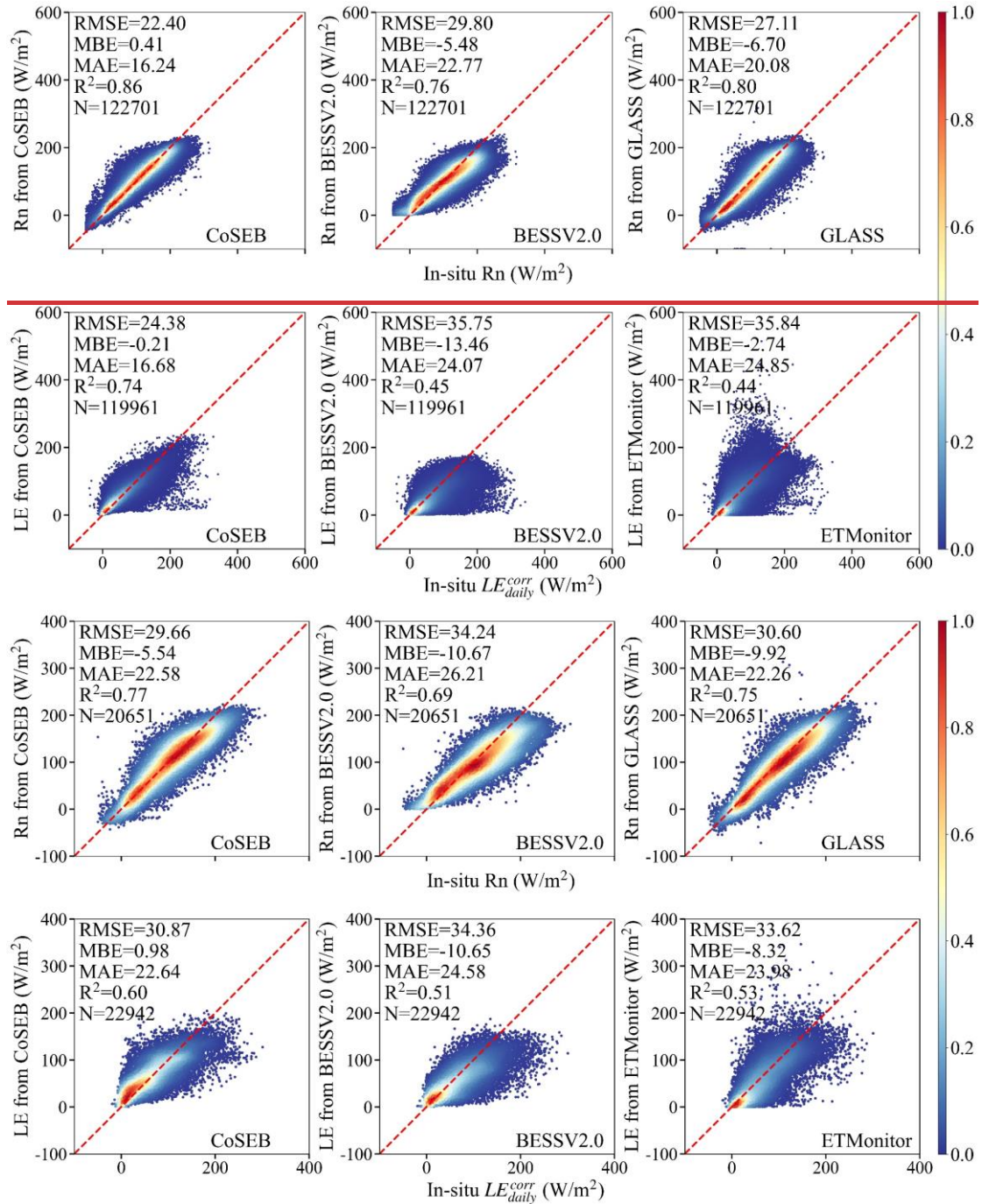
Fig. 6 and Fig. 7 present the comparison of daily  $SW_{IN}$ ,  $LW_{IN}$  and  $LW_{OUT}$ , as well as  $R_n$  and  $LE$  from the CoSEB-based datasets and mainstream products/datasets (including GLASS, BESS-Rad, BESSV2.0 and ETMonitor), with in situ observations at 134 EC test sites, respectively. Overall, the estimates from the CoSEB-based datasets exhibited a closer agreement with in situ observations than those from mainstream products/datasets, where the CoSEB-based datasets reduced the RMSE by  $4.35$  to  $0.01 \text{ W/m}^2$  and increased the  $R^2$  by  $0.04$  to  $0.09$  compared

to mainstream products. Specifically, the RMSE for the  $SW_{IN}$ ,  $LW_{IN}$ ,  $LW_{OUT}$  increased from ~~28.51~~37.52  $W/m^2$ , ~~14.29~~22.47  $W/m^2$  and ~~10.62~~13.78  $W/m^2$  in the CoSEB-based datasets to ~~35.44~~7.53  $W/m^2$ , ~~18.64~~23.37  $W/m^2$  and ~~15.29~~16.46  $W/m^2$  in the GLASS, respectively, and for  $SW_{IN}$  from ~~28.51~~37.52  $W/m^2$  in the CoSEB-based datasets to ~~36.23~~40.87  $W/m^2$  in the BESS-Rad. Likewise, the RMSEs for daily Rn and LE were ~~22.40~~9.66  $W/m^2$  and ~~24.38~~30.87  $W/m^2$  in the CoSEB-based datasets, which were lower than those of ~~29.80~~34.24  $W/m^2$  and ~~35.75~~4.36  $W/m^2$  in BESSV2.0, respectively, as well as those of ~~27.11~~30.60  $W/m^2$  for Rn in GLASS and ~~35.84~~3.62  $W/m^2$  for LE in ETMonitor.





**Fig. 6 Comparison of the daily downward shortwave radiation ( $SW_{IN}$ , the first column), downward longwave radiation ( $LW_{IN}$ , the second column) and upward longwave radiation ( $LW_{OUT}$ , the third column) from the CoSEB-based datasets, GLASS and BESS-Rad with the in situ observed  $SW_{IN}$ ,  $LW_{IN}$  and  $LW_{OUT}$  at 134-44 eddy-covariance test sites. The colorbar represents the normalized density of data points.**



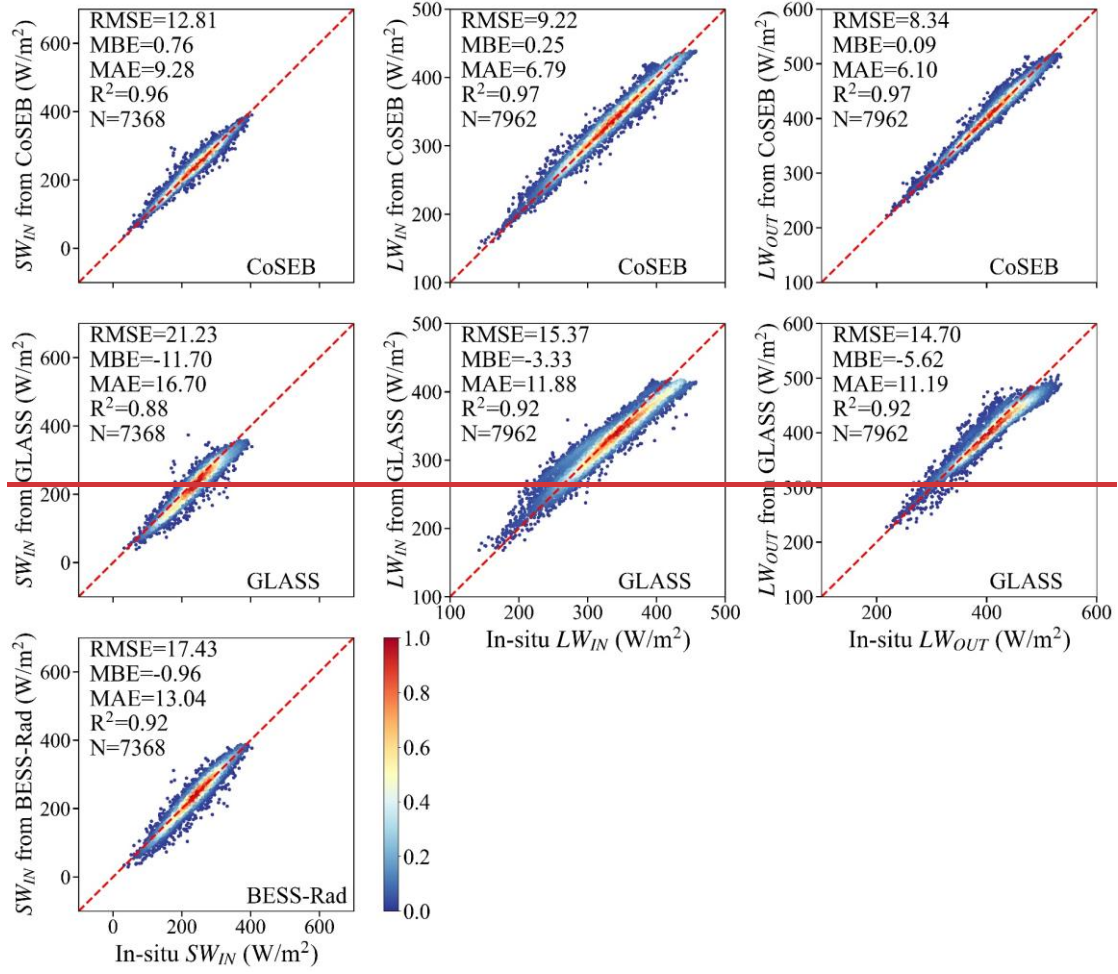
**Fig. 7 Comparison of the daily net radiation (Rn, the upper row) and latent heat flux (LE, the lower row) from the CoSEB-based datasets, BESSV2.0, GLASS and ETMonitor with the in situ observed Rn, and energy imbalance-corrected LE ( $LE_{daily}^{corr}$ ) at 134 eddy covariance test sites. The colorbar represents the normalized density of data points.**

Figs. 8, 9 and 10 compare the 8-day  $SW_{IN}$ ,  $LW_{IN}$  and  $LW_{OUT}$ , Rn and LE, as well as H from the CoSEB-based datasets and mainstream products, with in situ observations at 44 test sites, respectively. Overall, the CoSEB-based datasets outperformed



the mainstream products/datasets for all surface radiation and heat fluxes, where the CoSEB-based datasets reduced the RMSE by  $4.620.24$  W/m<sup>2</sup> to  $14.640.48$  W/m<sup>2</sup> and increased the R<sup>2</sup> by  $0.0401$  to  $0.4138$  compared to mainstream products. Specifically, for  $SW_{IN}$ ,  $LW_{IN}$  and  $LW_{OUT}$ , the RMSE increased from  $12.818.54$  W/m<sup>2</sup>,  $9.2218.50$  W/m<sup>2</sup> and  $8.349.41$  W/m<sup>2</sup> in the CoSEB-based datasets to  $21.2335$  W/m<sup>2</sup>,  $15.3720.39$  W/m<sup>2</sup> and  $14.7048$  W/m<sup>2</sup> in the GLASS, respectively, and for  $SW_{IN}$  from  $12.818.54$  W/m<sup>2</sup> in the CoSEB-based datasets to  $17.4318.78$  W/m<sup>2</sup> in the BESS-Rad. For R<sub>n</sub>, the RMSE increased from  $13.389.12$  W/m<sup>2</sup> in the CoSEB-based datasets to  $\sim 23$  W/m<sup>2</sup> in the FLUXCOM and GLASS and to  $>27$  W/m<sup>2</sup> in the BESSV2.0 and  $>23$  W/m<sup>2</sup> in the FLUXCOM and BESSV2.0, while the R<sup>2</sup> decreased from  $0.9182$  in the CoSEB-based datasets to  $0.75$  in the FLUXCOM and GLASS and to  $0.8262$  in the GLASS-BESSV2.0 and to  $<0.72$  in the FLUXCOM and BESSV2.0. Likewise, for LE, the RMSE increased from  $19.9922.31$  W/m<sup>2</sup> in the CoSEB-based datasets to  $\sim 26.1625$  W/m<sup>2</sup> in the FLUXCOM, PML V2, BESSV2.0 and ETMonitor, and to  $>28.1732$  W/m<sup>2</sup> in BESSV2.0, MOD16A2, PML\_V2 and ETMonitor, while the R<sup>2</sup> decreased from  $0.867$  in the CoSEB-based datasets to  $\sim 0.6560$  in the FLUXCOM, PML\_V2, BESSV2.0 and ETMonitor, and to  $<0.63$  in the remaining products MOD16A1. For H, the RMSE increased from  $17.4421.63$  W/m<sup>2</sup> in the CoSEB-based datasets to  $23.962.64$  W/m<sup>2</sup> in the FLUXCOM.

The differences between the estimates from the CoSEB-based datasets and mainstream datasets are likely multifactorial, arising from the simplification and parameterization uncertainties in physics-based models, as well as the lack of physical constraints, limited training samples, and incomplete consideration of influencing factors in other machine-learning-based models.





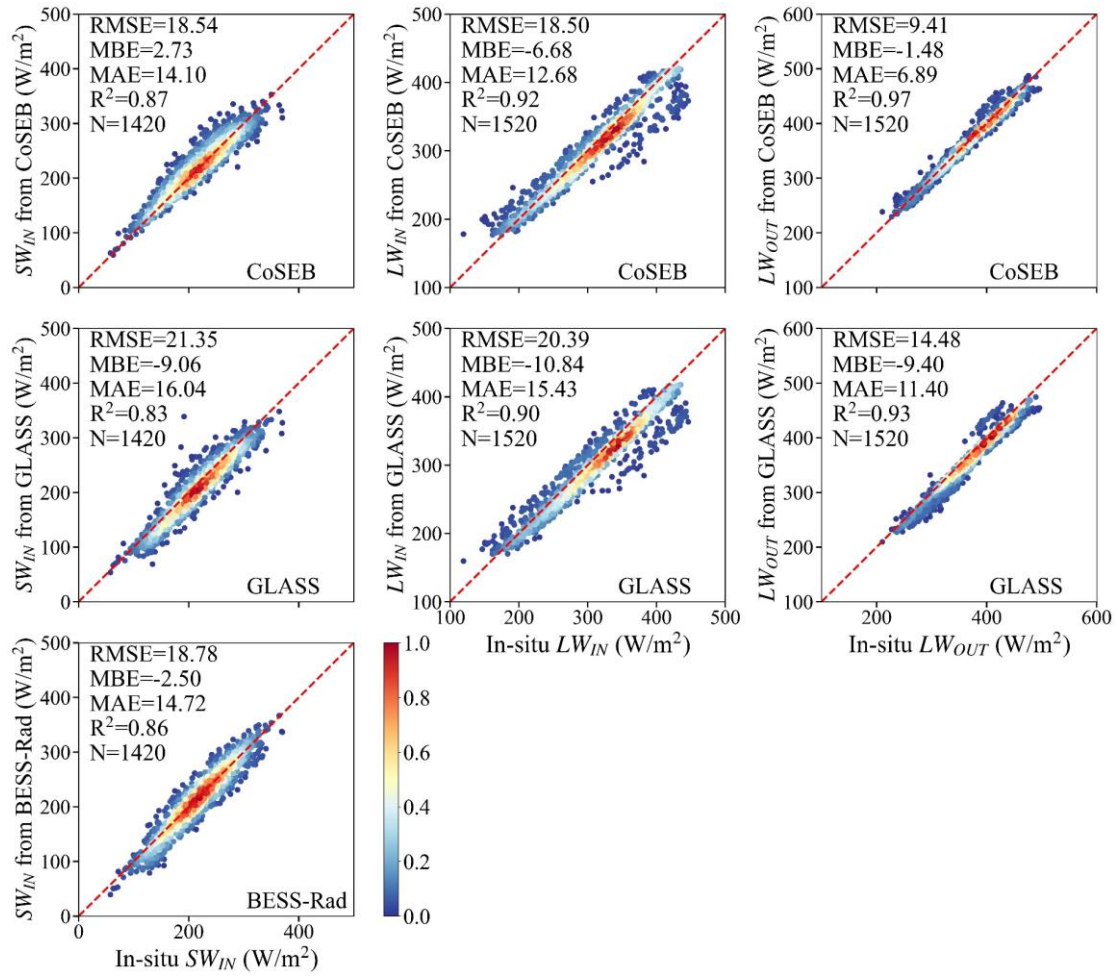
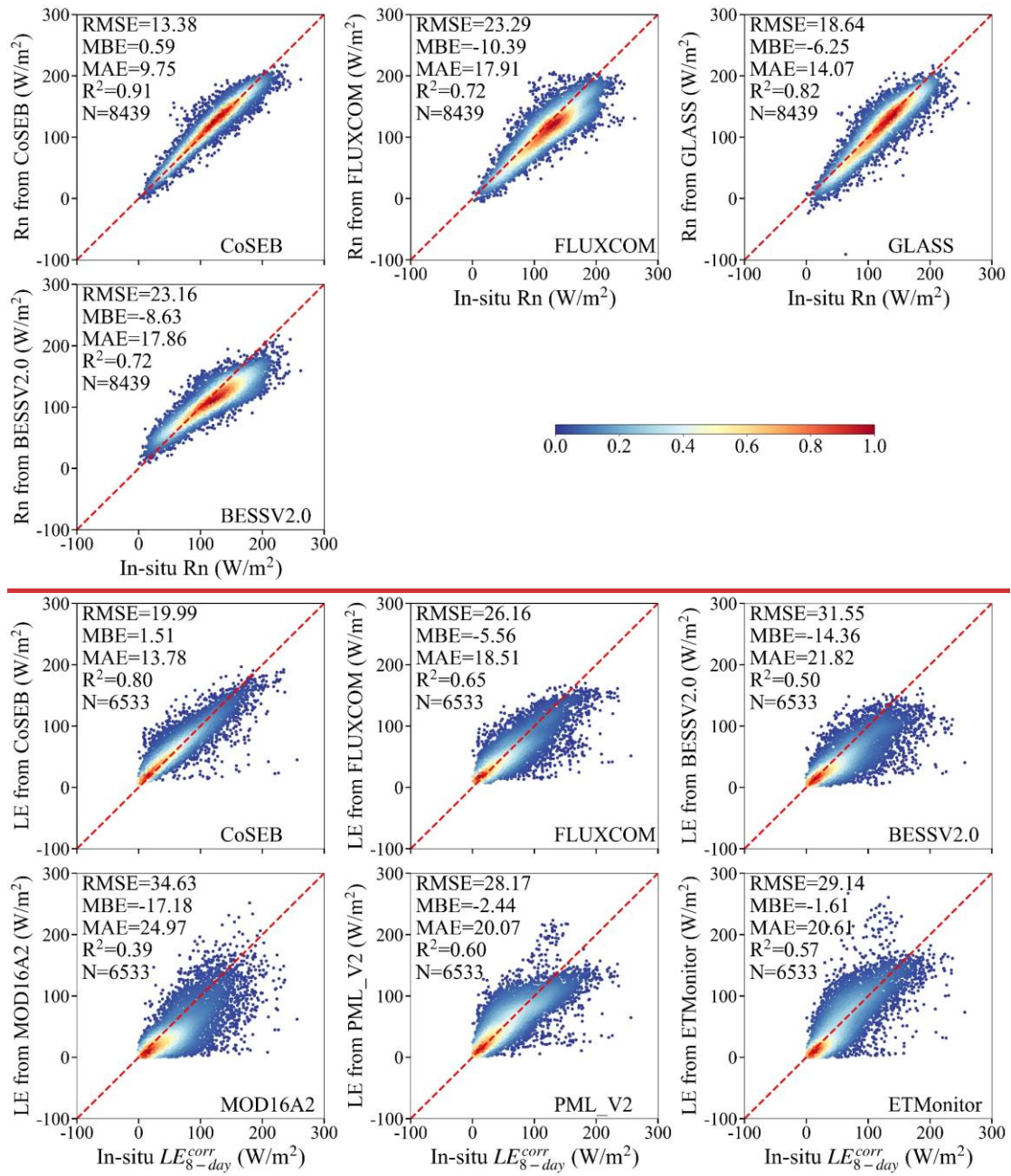
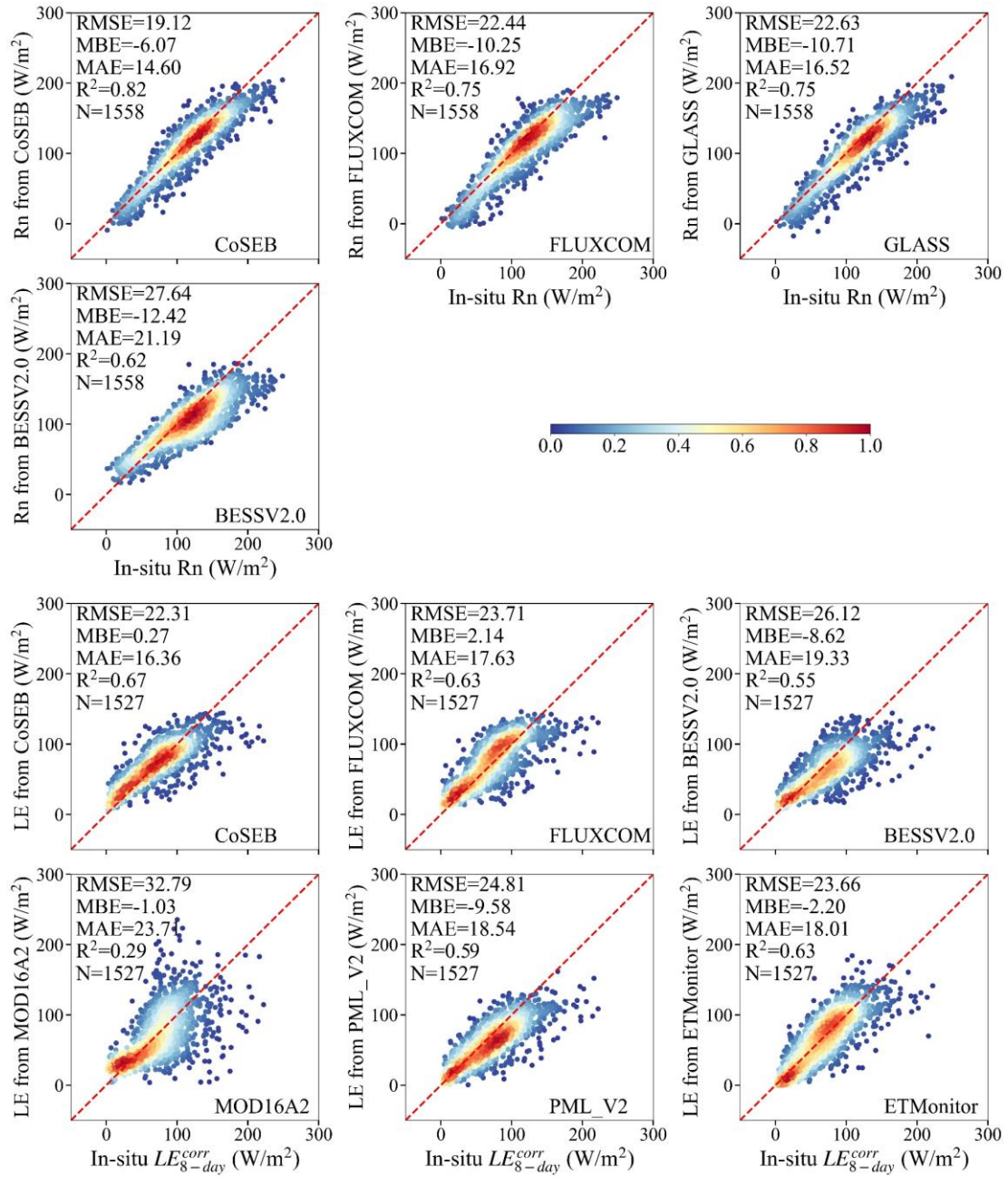


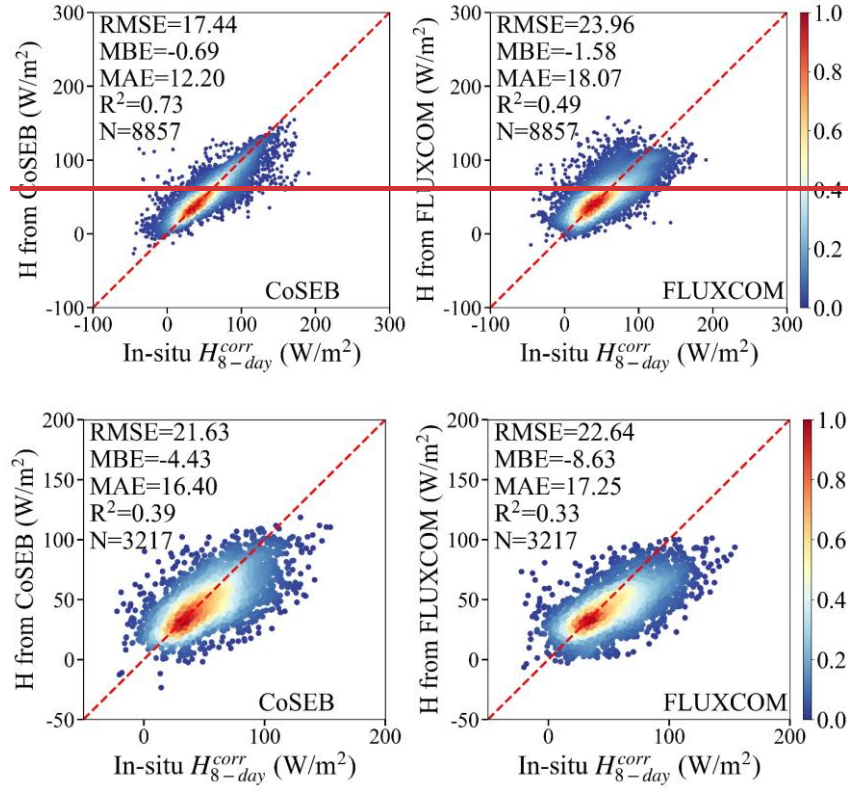
Fig. 8 Same as Fig. 6, but for the comparison at 8-day scale.



504



**Fig. 9 Comparison of the 8-day net radiation (Rn, the upper two rows) and latent heat flux (LE, the lower three rows) from the CoSEB-based datasets, FLUXCOM, BESSV2.0, GLASS, MOD16A2, PML\_V2 and ETMonitor with in situ observed Rn, and energy imbalance-corrected LE ( $LE_{8-day}^{corr}$ ) at 134-44 testddy-covariance sites. The colorbar represents the normalized density of data points.**



**Fig. 10** Comparison of the 8-day sensible heat flux ( $H$ ) from the CoSEB-based datasets and the FLUXCOM with the in situ energy imbalance-corrected  $H$  ( $H_{8-day}^{corr}$ ) at 134-44 eddy covariance sites. The colorbar represents the normalized density of data points.

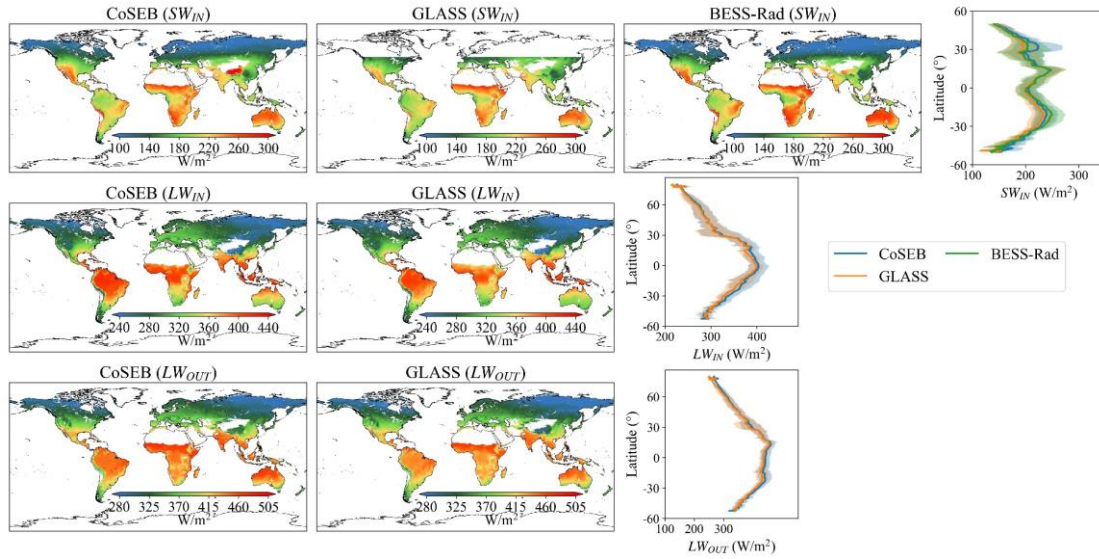
### 4.3 Spatial-temporal patterns of global land surface radiation and heat fluxes

In addition to the validation and inter-comparison of the CoSEB-based datasets at the global-site scales, we further inter-compared the estimates of land surface radiation and heat fluxes from the CoSEB-based datasets and the mainstream products/datasets, in terms of their global spatial and temporal patterns.

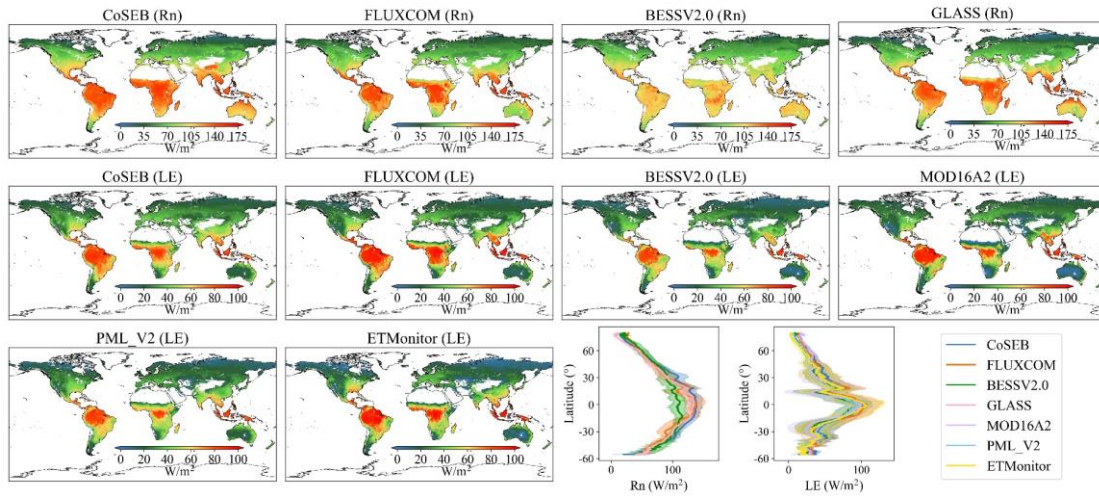
Figs. 11, 12 and 13 show the spatial distributions (excluding Greenland, Antarctic continent, deserts, water bodies and permanent snow) and latitudinal profiles of the global  $0.05^\circ$  mean annual  $SW_{IN}$ ,  $LW_{IN}$  and  $LW_{OUT}$ ,  $R_n$  and  $LE$ , as well as  $H$  from 2001 to 2018, respectively, as derived from the CoSEB-based datasets and mainstream products/datasets [i.e. GLASS, BESS-Rad, BESSV2.0, FLUXCOM, MOD16A2, PML\_V2 and ETMonitor, resampled to  $0.05^\circ$  using arithmetic averaging method or cubic convolutional method if necessary]. Overall, the spatial patterns of the estimates

529 from the CoSEB-based datasets aligned well with those observed in these mainstream  
530 products/datasets, though regional discrepancies were present. Specifically, the mean  
531 annual  $LW_{IN}$ ,  $LW_{OUT}$ ,  $R_n$ , and  $LE$  generally exhibited decreasing trends from the equator  
532 towards higher latitudes, peaking in regions such as the Amazon Rainforest, Congo  
533 Rainforest, and the Malay Archipelago. In contrast, the higher mean annual  $SW_{IN}$  and  
534  $H$  were mainly found in the Tibetan Plateau, southwestern U.S., mid-west Australia,  
535 Sahel and Southern Africa, while the lower values were found in high-latitude regions  
536 of  $>50^\circ N$ . In the region ~~with~~ of high values, the mean annual estimates of  $SW_{IN}$  from  
537 the CoSEB-based datasets were higher than those from GLASS but lower than those  
538 from BESS-Rad, the estimates of  $LW_{IN}$  and  $LW_{OUT}$  from the CoSEB-based datasets were  
539 both higher than those from GLASS, the estimates of  $R_n$  from the CoSEB-based  
540 datasets were significantly higher than those from BESSV2.0, and comparable to or  
541 slightly higher than those from FLUXCOM and GLASS, the estimates of  $LE$  from the  
542 CoSEB-based datasets were close to those from BESSV2.0 and PML\_V2, but slightly  
543 lower than those from FLUXCOM, MOD16A2 and ETMonitor. Besides, the estimates  
544 of  $H$  from the CoSEB-based datasets were higher than those from FLUXCOM in  
545 regions with high values, while lower than those from FLUXCOM in regions with low  
546 values.



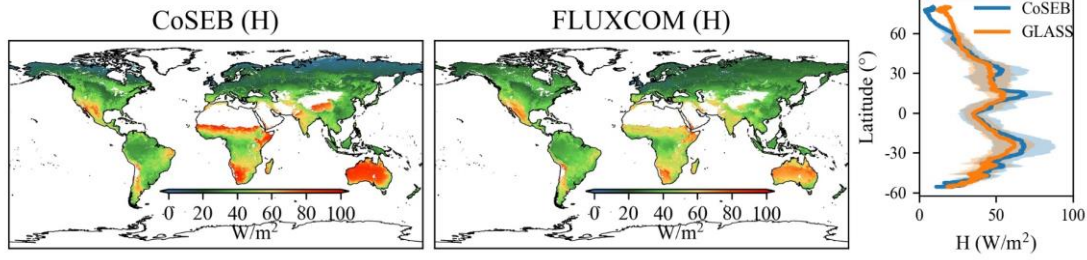


**Fig.11 Spatial patterns of global mean annual downward shortwave radiation ( $SW_{IN}$ , the first row), downward longwave radiation ( $LW_{IN}$ , the second row) and upward longwave radiation ( $LW_{OUT}$ , the third row) from 2001 to 2018 by CoSEB-based datasets, GLASS and BESS-Rad. The rightmost subfigure of each row represents the latitudinal profiles of mean annual  $SW_{IN}$ ,  $LW_{IN}$  and  $LW_{OUT}$  from CoSEB-based datasets, GLASS and BESS-Rad, where the shaded area represents the variation of standard deviation for each product.**



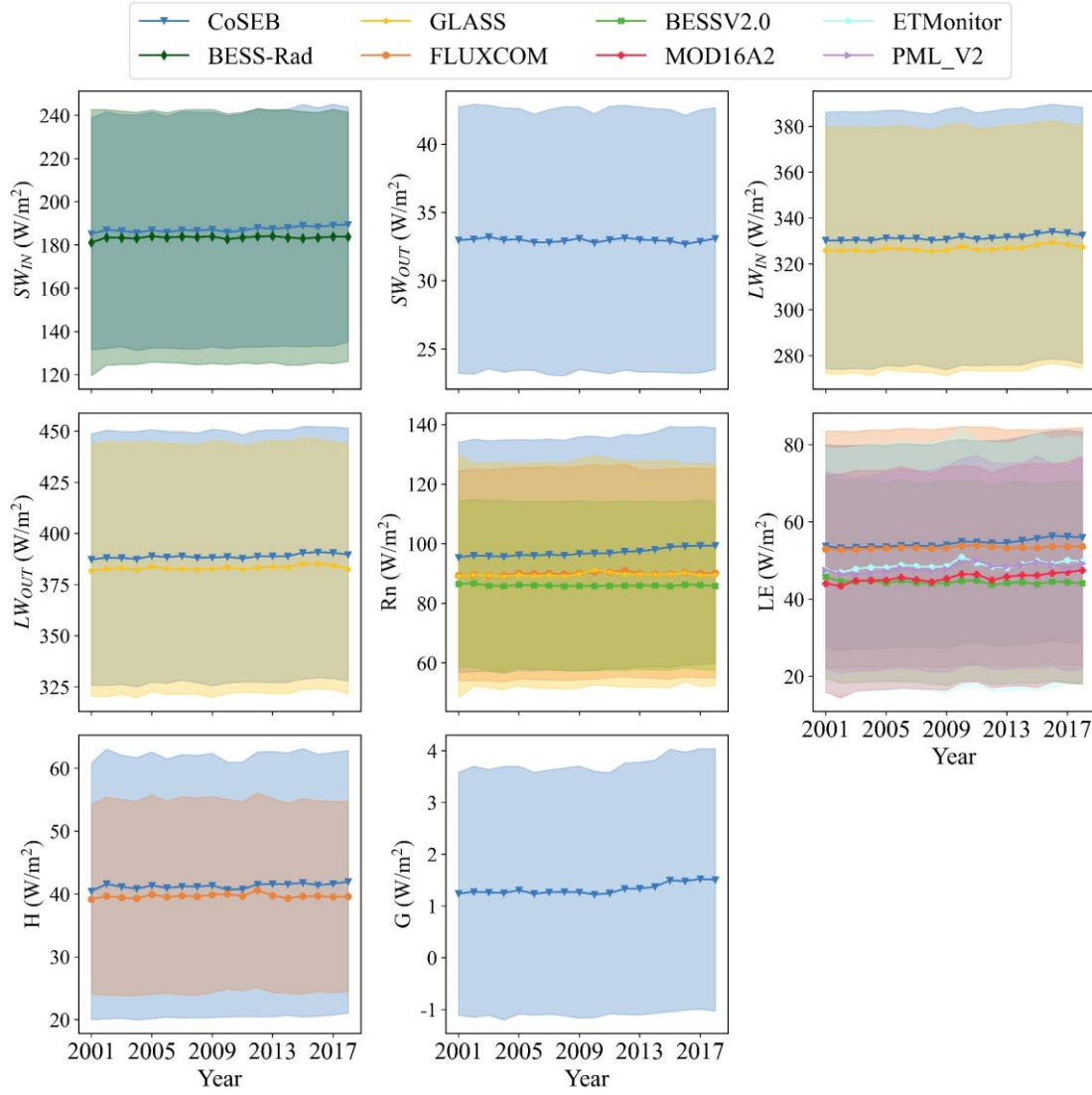
**Fig.12 Spatial patterns of global mean annual net radiation ( $Rn$ , the first row) and latent heat flux ( $LE$ , the second and third rows) from 2001 to 2018 by CoSEB-based datasets, FLUXCOM, BESSV2.0, MOD16A2, PML\_V2, ETMonitor and GLASS. The last two subfigures of the third row represent the latitudinal profiles of mean annual  $Rn$  and  $LE$  from CoSEB-based datasets and these mainstream products/datasets, where the shaded area represents the variation of standard deviation for each product.**



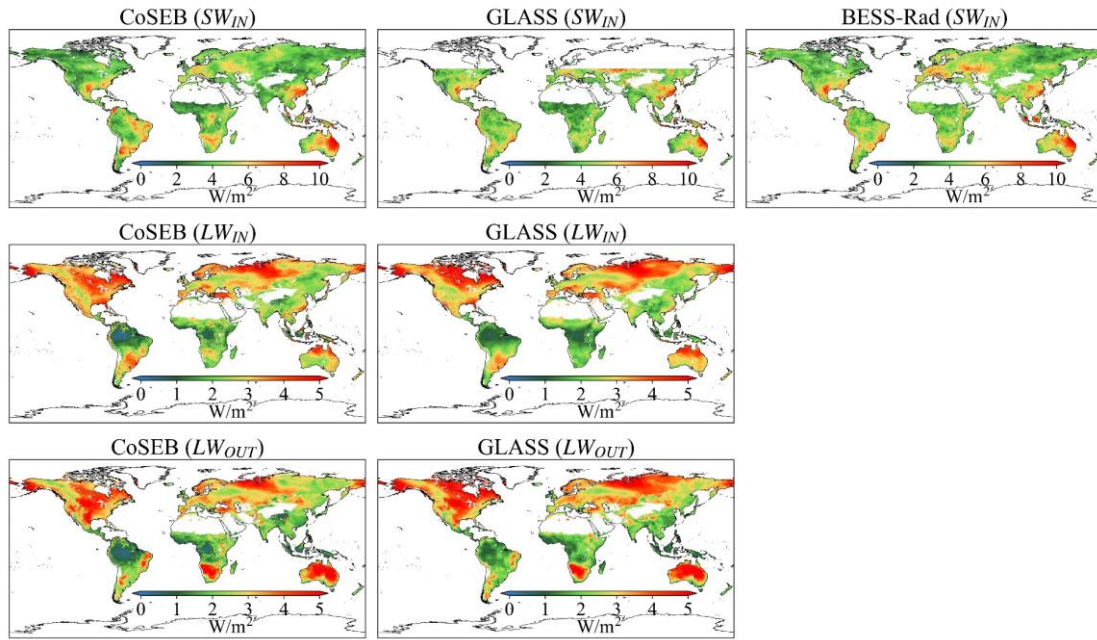


**Fig.13 Spatial patterns of global mean annual sensible heat flux ( $H$ ) from 2001 to 2018 by CoSEB-based datasets and FLUXCOM. The rightmost subfigure represents the latitudinal profiles of mean annual  $H$  from CoSEB-based datasets and FLUXCOM, where the shaded area represents the variation of standard deviation for each product.**

The temporal evolutions of the global (excluding Greenland, Antarctic continent, deserts, water bodies and permanent snow) land surface radiation and heat fluxes derived from the CoSEB-based datasets and mainstream products/datasets from 2001 to 2018 were also investigated, as shown in Fig. 14. The results indicated that the temporal variation of each flux from the CoSEB-based datasets generally agreed well with those from mainstream products/datasets, exhibiting relatively stable trends. The global annual mean estimates using area weighting average by the CoSEB-based datasets from 2001 to 2018 varied between  $\sim 185.22$  and  $\sim 189.50$   $\text{W/m}^2$  with the mean of  $\sim 187.23$   $\text{W/m}^2$  for  $SW_{IN}$ , between  $\sim 32.67$  and  $\sim 33.20$   $\text{W/m}^2$  with the mean of  $\sim 32.96$   $\text{W/m}^2$  for  $SW_{OUT}$ , between  $\sim 330.24$  and  $\sim 334.14$   $\text{W/m}^2$  with the mean of  $\sim 331.50$   $\text{W/m}^2$  for  $LW_{IN}$ , between  $\sim 387.25$  and  $\sim 390.82$   $\text{W/m}^2$  with the mean of  $\sim 388.81$   $\text{W/m}^2$  for  $LW_{OUT}$ , between  $\sim 95.41$  and  $\sim 99.39$   $\text{W/m}^2$  with the mean of  $97.11$   $\text{W/m}^2$  for  $R_n$ , between  $\sim 53.24$  and  $\sim 56.37$   $\text{W/m}^2$  with the mean of  $\sim 54.53$   $\text{W/m}^2$  for  $LE$ , between  $\sim 40.44$  and  $\sim 41.96$   $\text{W/m}^2$  with the mean of  $\sim 41.29$   $\text{W/m}^2$  for  $H$ , and between  $\sim 1.22$  and  $\sim 1.52$   $\text{W/m}^2$  with the mean of  $\sim 1.33$   $\text{W/m}^2$  for  $G$ . For each radiation or heat flux, the annual mean estimates from the CoSEB-based datasets were overall higher than those from the mainstream products/datasets. In particular, the annual mean  $R_n$  estimates from the CoSEB-based datasets were higher than those from FLUXCOM, GLASS and BESSV2.0 sequentially, and the annual mean  $LE$  estimates from the CoSEB-based datasets were marginally higher than those from FLUXCOM, but substantially exceeded those from ETMonitor, PML\_V2, MOD16A2 and BESSV2.0 sequentially.



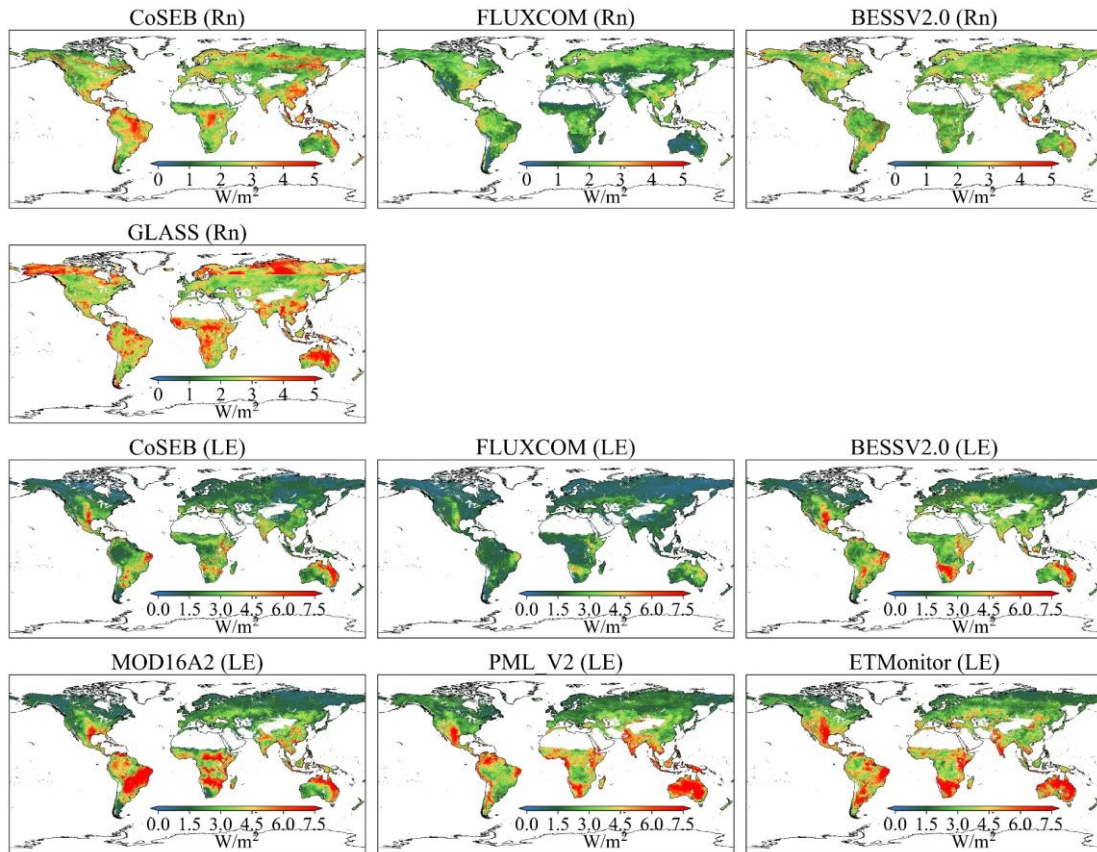
**Fig. 14** Temporal variation of annual mean downward shortwave radiation ( $SW_{IN}$ ), upward shortwave radiation ( $SW_{OUT}$ ), downward longwave radiation ( $LW_{IN}$ ), upward longwave radiation ( $LW_{OUT}$ ), net radiation (Rn), latent heat flux (LE), sensible heat flux (H) and soil heat flux (G) from 2001 to 2018 from the CoSEB-based datasets, BESS-Rad, GLASS, FLUXCOM, BESSV2.0, PML\_V2, MOD16A2 and ETMonitor. The shaded area represents the variation of standard deviation for each product.



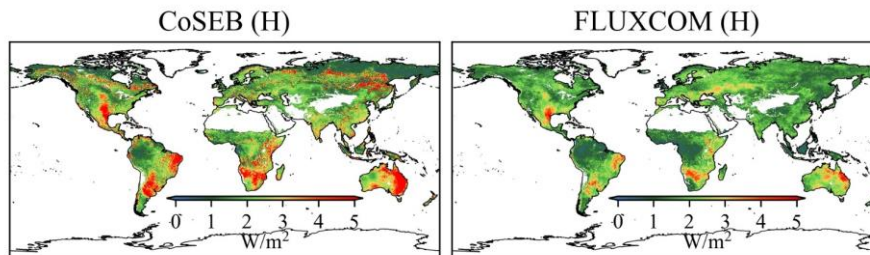
**Fig. 15** Spatial distribution of interannual variability (standard deviation) of downward shortwave radiation ( $SW_{IN}$ , the first row), downward longwave radiation ( $LW_{IN}$ , the second row) and upward longwave radiation ( $LW_{OUT}$ , the third row) from 2001 to 2018 by the CoSEB-based datasets, GLASS and BESS-Rad.

Figs. 15, 16 and 17 show the spatial patterns (excluding Greenland, Antarctic continent, deserts, water bodies and permanent snow) of interannual variability of  $SW_{IN}$ ,  $LW_{IN}$  and  $LW_{OUT}$ ,  $R_n$  and  $LE$ , as well as  $H$  from 2001 to 2018, respectively, derived from the CoSEB-based datasets and mainstream products/datasets. In general, the estimates from the CoSEB-based datasets displayed similar interannual variability in space with those from the mainstream products/datasets. Specially, the estimates of  $SW_{IN}$  from the CoSEB-based datasets, BESS-Rad, and GLASS exhibited a significant interannual variability mainly in northeastern Australia, eastern South America, Southeast China, and Southwest North America. The interannual variability of  $LW_{IN}$  and  $LW_{OUT}$  by the CoSEB-based datasets and GLASS displayed high values primarily at middle-to-high latitudes of the Northern Hemisphere and parts of Africa and Australia. The interannual variability of  $R_n$  observed by the CoSEB-based datasets was generally lower than that of GLASS, but higher than that of BESSV2.0 and FLUXCOM. The CoSEB-based datasets missed the strong interannual variability of  $LE$  as observed in MOD16A2, PML\_V2 and ETMonitor in parts of Africa, Australia and eastern South

America. Furthermore, FLUXCOM exhibited the weakest interannual variability of LE in almost all regions. The interannual variability of H derived from the CoSEB-based datasets was higher than ~~those~~ that from FLUXCOM, with stronger interannual variabilities mainly observed in parts of eastern South America, southern Africa, and northeastern Australia.



**Fig. 16** Spatial distribution of interannual variability (standard deviation) of net radiation (Rn, the first and second rows) and latent heat flux (LE, the third and fourth row) from 2001 to 2018 by the CoSEB-based datasets, FLUXCOM, BESSV2.0, MOD16A2, PML\_V2, ETMonitor and GLASS.



**Fig. 17** Spatial distribution of interannual variability (standard deviation) of sensible heat flux (H) from 2001 to 2018 by the CoSEB-based datasets and FLUXCOM.



## 5 Discussion

Accurately monitoring the spatial and temporal variations of global land surface radiation and heat fluxes is crucial for quantifying the exchange of radiation, heat and water between the land and atmosphere under global climate change (Chen et al., 2020; Du et al., 2024; Kim et al., 2023; Liang et al., 2006; Wang et al., 2020). However, although numerous global RS-based products/datasets of land surface radiation and heat fluxes have been developed using physical and/or statistical methods, they typically provide either merely a single flux or multiple fluxes (see Table 1) that are estimated separately from uncoordinated models (Huang et al., 2024; Jung et al., 2019; Sun et al., 2023; Tang et al., 2019), leading to noticeable radiation imbalance and/or heat imbalance when these products are combined for practical applications. To address these limitations, we generated high-accuracy global datasets of land surface radiation and heat fluxes from 2000 to 2020 that adhere to both radiation and heat conservation laws, using our proposed CoSEB model (Wang et al., 2025).

Our CoSEB model, integrating underlying physical principles of training datasets into machine learning technique to effectively learn the interrelations among multiple targeted outputs, was originally designed for coordinating estimates of global land surface energy balance components ( $R_n$ ,  $LE$ ,  $H$  and  $G$ ) to satisfy the energy conservation (Wang et al., 2025). Inspired by the idea of constructing the original CoSEB model, we further incorporated land surface radiation fluxes into our model to simultaneously consider the physical constraints of both surface radiation and heat conservation principles, by renewing the CoSEB using multiple remote sensing ~~products and~~ reanalysis datasets, as well as in-situ observations of  $SW_{IN}$ ,  $SW_{OUT}$ ,  $LW_{IN}$ ,  $LW_{OUT}$ ,  $R_n$ ,  $LE$ ,  $H$  and  $G$ . In selecting the 19 input variables to accommodate the additional target variables, prior knowledge derived from previous studies was employed to identify factors that exert significant influence on surface radiation and heat flux while maintaining relative inter-independence as much as possible (Jung et al., 2019; Mohan et al., 2020; Wang et al., 2021; Xian et al., 2024). This practice is

commonly adopted in data-driven models for estimating land surface water, energy, and carbon fluxes (Bai et al., 2024; Elghawi et al., 2023; Han et al., 2023; O. & Orth, 2021). The importance scores of the 19 different feature variables are exhibited in Table S4 in the Supplementary Material, and downward solar radiation, the primary source of the energy at the earth surface, is the most important input variable, consistent with the results from our previous study (Wang et al., 2025). Although some of the selected variables may exhibit a certain degree of multi-collinearity, each contributes unique and physically meaningful information, supporting the inclusion of all variables in model construction. To comprehensively account for the main factors influencing surface radiation and heat fluxes (Mohan et al., 2020; Wang et al., 2021; Xian et al., 2024) [JW1], the renewed CoSEB model utilized 19 easily accessible parameters/variables from ERA5 Land reanalysis datasets, GLASS products, MODIS products, GMTED2010 and NOAA/GML as input, which were readily available to generate datasets of global land surface radiation and heat fluxes in a practical and operational manner. (Wang et al., 2025) Note that the variable importance, derived from the built-in method of the random forests and potentially affected by multicollinearity among the input variables, is presented only as a reference. Retaining all 19 feature variables ensures the model's flexibility and generalization capability, enabling future incorporation of additional representative ground-based observations for further training and improvement. Besides, to investigate the impact of lagged effects of input variables on model performance, experiments were also conducted by adding lagged variables (e.g., the air temperature of the previous day) to the 19 input features. The results (Fig. S4 in the Supplementary Material) showed almost no improvement in model accuracy, suggesting that lagged effects on model performance were negligible within the CoSEB framework for estimates of daily surface radiation and heat fluxes. Furthermore, to better illustrate the effect of including additional radiation components ( $SW_{IN}$ ,  $SW_{OUT}$ ,  $LW_{IN}$  and  $LW_{OUT}$ ) in the renewed CoSEB model compared with the original version by Wang et al. (2025), we have tested the performance of a reconstructed model that



estimated only  $R_n$ ,  $LE$ ,  $H$  and  $G$  using the same independent variables and samples as those in the renewed CoSEB model. The results (Fig. S5 in the supplementary material) showed no significant differences in accuracy compared with those of the renewed CoSEB model, indicating the expansion of radiation components did not compromise model performance.

The main advantages of our CoSEB-based datasets of land surface radiation and heat fluxes lie in that [1] they are the first RS-based data-driven global datasets that satisfy both surface radiation balance ( $SW_{IN} - SW_{OUT} + LW_{IN} - LW_{OUT} = R_n$ ) and heat balance ( $LE + H + G = R_n$ ) among the eight fluxes, as demonstrated by both the RIR and EIR of 0, [2] the radiation and heat fluxes are characterized by high accuracies when validated against in-situ measurements at 134 “homogeneous” 44 independent test sites (see the first-second paragraph in Section 4.2.2.1), where (1) the RMSEs for daily estimates of  $SW_{IN}$ ,  $SW_{OUT}$ ,  $LW_{IN}$ ,  $LW_{OUT}$ ,  $R_n$ ,  $LE$ ,  $H$  and  $G$  from the CoSEB-based datasets were 28.5137.52 W/m<sup>2</sup>, 10.394.20 W/m<sup>2</sup>, 14.2922.47 W/m<sup>2</sup>, 10.623.78 W/m<sup>2</sup>, 22.409.66 W/m<sup>2</sup>, 24.3830.87 W/m<sup>2</sup>, 22.679.75 W/m<sup>2</sup> and 6.775.69 W/m<sup>2</sup>, respectively, as well as for 8-day estimates were 12.848.54 W/m<sup>2</sup>, 7.0812.19 W/m<sup>2</sup>, 9.2218.50 W/m<sup>2</sup>, 8.349.41 W/m<sup>2</sup>, 13.389.12 W/m<sup>2</sup>, 19.9922.31 W/m<sup>2</sup>, 17.4421.63 W/m<sup>2</sup> and 4.254.60 W/m<sup>2</sup>, respectively, (2) the CoSEB-based datasets, in comparison to the mainstream RS-based products/datasets (i.e. GLASS, BESS-Rad, FLUXCOM, BESSV2.0, MOD16A2, PML\_V2 and ETMonitor), better agreed with the in situ observations at 134 EC the 44 test sites, showing the RMSE reductions ranging from 4.350.01 W/m<sup>2</sup> to 11.464.58 W/m<sup>2</sup> for  $SW_{IN}$ ,  $LW_{IN}$ ,  $LW_{OUT}$ ,  $R_n$  and  $LE$  at daily scale, and 4.620.24 W/m<sup>2</sup> to 14.640.48 W/m<sup>2</sup> for  $SW_{IN}$ ,  $LW_{IN}$ ,  $LW_{OUT}$ ,  $R_n$ ,  $LE$  and  $H$  at 8-day scale. Furthermore, the CoSEB-based datasets outperformed the ERA5-Land reanalysis datasets in estimating surface energy fluxes (where  $SW_{OUT}$ ,  $LW_{OUT}$ ,  $R_n$  and  $G$  for the ERA-Land were inferred from surface radiation balance and heat balance), particularly for  $SW_{OUT}$ ,  $H$  and  $G$ , with RMSE reductions of 0.13-8.15 W/m<sup>2</sup> when validated against in situ

observations at the 44 test sites (Figs. S6 and S7 in the Supplementary Material). Preliminary analysis indicates that the CoSEB-based datasets exhibit spatial patterns consistent with those of mainstream RS-based datasets and Earth system model outputs (see Fig. S8 in the supplementary material). More detailed analysis about their similarities and differences can be further conducted in future work.

Our developed datasets could be potentially applied in many fields, including but not limited to (1) exploring the spatial-temporal patterns of global land surface radiation and heat flux (es) and their driving mechanisms over the past decades under global change (e.g., rising CO<sub>2</sub> concentration, greening land surface and increasing air temperature), (2) investigating the variability of land surface radiation and heat fluxes caused by extreme events and human activities, e.g. afforestation or deforestation, wildfire, air pollution, weather extremes and urbanization, (3) assessing the resources of solar energy, geothermal energy, surface and ground water at regional and global scales, (4) monitoring natural hazards, e.g. drought in agriculture and forestry.

The uncertainties of our datasets are relevant to (1) the data preprocessing, and (2) the application of the CoSEB ~~at-model across~~ different spatial scales. Specifically, ~~the~~ daily averages of surface radiation and heat fluxes for each day ~~were~~ obtained for analysis from good-quality half-hourly observations when the fraction of these good-quality half-hourly observations was greater than 80% in a day, due to the lack of consensus on the method for aggregating gapped half-hourly observations to daily data (Tang et al., 2024a; Yao et al., 2017; Zheng et al., 2022). Simple temporal interpolation of half-hourly in situ observations, which could therefore introduce substantial uncertainties, was not applied, because surface radiation and heat fluxes are sensitive to short-term variations in meteorological conditions and their intraday dynamics are often complex. Likewise, since there was no agreement on how to correct for the energy imbalance of turbulent heat fluxes, we adopted the most widely applied Bowen ratio method to enforce energy closure between  $R_n - G$  and  $LE + H$  (Castelli et al., 2018; Twine et al., 2000; Zhang et al., 2021). Another potential source of uncertainty arises

from differences in meteorological reanalysis data caused by spatial downscaling, which, as demonstrated in our previous study (Wang et al., 2025, the last paragraph of Section 5.1), has a relatively small impact on model estimates by the machine-learning-based CoSEB model combined with finer-resolution surface-related variables that partially compensate for the spatial heterogeneity and localized variations not captured by the coarse-resolution datasets. ~~(Wang et al., 2025, the last paragraph of Section 5.1)~~ ~~(Wang et al., 2025, the last paragraph of Section 5.1)~~ These data preprocessing had an effect on the construction of the renewed CoSEB model, which may further affect the global datasets. Moreover, the renewed CoSEB model was constructed at the spatial scale of 500 m to match the footprints of the in situ EC observations, but applied at the spatial resolution of 0.05° to generate global datasets, mainly limited by the computing and storage capabilities ~~in~~of our personal computers. However, the CoSEB-based datasets have also been validated and inter-compared at ~~134 EC~~44 independent test sites to demonstrate that the difference in spatial scale would not much affect the performance of the datasets. Despite these uncertainties, it is worth emphasizing that our work was the first attempt to innovatively develop data-driven energy-conservation datasets of global land surface radiation and heat fluxes with high accuracies.

## 6 Data availability

The energy-conservation datasets of global land surface radiation and heat fluxes generated by the CoSEB model with spatial-temporal resolutions of daily and 0.05° from Feb.26, 2000 to Dec.31, 2020 are freely available through the National Tibetan Plateau Data Center at <https://doi.org/10.11888/Terre.tpd.c.302559> (Tang et al., 2025a) and through the Science Data Bank (ScienceDB) at <https://doi.org/10.57760/sciencedb.27228> (Tang et al., 2025b).

## 7 Summary and Conclusion

This study for the first time developed data-driven energy-conservation datasets of global land surface radiation and heat fluxes using our CoSEB model renewed based on GLASS and MODIS products, ERA5-Land reanalysis datasets, topographic data, CO<sub>2</sub> concentration data, and observations at 258 EC sites worldwide ~~from the FLUXNET, AmeriFlux, EuroFlux, OzFlux, ChinaFLUX and TPDC.~~

The CoSEB-based datasets of land surface radiation and heat fluxes are the first ~~RS-based~~ data-driven global datasets that satisfy both surface radiation balance ( $\overline{SW_{IN}} - \overline{SW_{OUT}} + \overline{LW_{IN}} - \overline{LW_{OUT}} = R_n$ ) and heat balance ( $\overline{LE} + \overline{H} + \overline{G} = R_n$ ) among the eight fluxes. Meanwhile, the CoSEB-based datasets outperformed the mainstream products/datasets in accuracy. Specifically, at ~~134-44 EC independent test sites, the RMSEs (R<sup>2</sup>) for daily estimates of  $\overline{SW_{IN}}$ ,  $\overline{SW_{OUT}}$ ,  $\overline{LW_{IN}}$ ,  $\overline{LW_{OUT}}$ ,  $\overline{R_n}$ ,  $\overline{LE}$ ,  $\overline{H}$  and  $\overline{G}$  from the CoSEB-based datasets were 37.52 W/m<sup>2</sup> (0.81), 14.20 W/m<sup>2</sup> (0.42), 22.47 W/m<sup>2</sup> (0.90), 13.78 W/m<sup>2</sup> (0.95), 29.66 W/m<sup>2</sup> (0.77), 30.87 W/m<sup>2</sup> (0.60), 29.75 W/m<sup>2</sup> (0.44) and 5.69 W/m<sup>2</sup> (0.44), respectively~~ the RMSEs for daily estimates of  $\overline{SW_{IN}}$ ,  $\overline{SW_{OUT}}$ ,  $\overline{LW_{IN}}$ ,  $\overline{LW_{OUT}}$ ,  $\overline{R_n}$ ,  $\overline{LE}$ ,  $\overline{H}$  and  $\overline{G}$  from the CoSEB-based datasets were 28.51 W/m<sup>2</sup>, 10.39 W/m<sup>2</sup>, 14.29 W/m<sup>2</sup>, 10.62 W/m<sup>2</sup>, 22.40 W/m<sup>2</sup>, 24.38 W/m<sup>2</sup>, 22.67 W/m<sup>2</sup> and 6.77 W/m<sup>2</sup>, respectively, as well as for 8-day estimates were ~~12.84~~ 18.54 W/m<sup>2</sup> (0.87), ~~7.08~~ 12.19 W/m<sup>2</sup> (0.39), ~~9.22~~ 18.50 W/m<sup>2</sup> (0.92), ~~8.34~~ 9.41 W/m<sup>2</sup> (0.97), ~~13.38~~ 9.12 W/m<sup>2</sup> (0.82), ~~19.99~~ 22.31 W/m<sup>2</sup> (0.67), ~~17.44~~ 21.63 W/m<sup>2</sup> (0.39) and ~~4.25~~ 4.60 W/m<sup>2</sup> (0.47), respectively. Moreover, the estimates from the CoSEB-based datasets in comparison to those from the mainstream products/datasets reduced the RMSE by ~~4.35~~ 0.01 W/m<sup>2</sup> to ~~11.46~~ 4.58 W/m<sup>2</sup> and increased the R<sup>2</sup> by ~~0.04~~ 0.01 to ~~0.30~~ 0.09 for  $\overline{SW_{IN}}$ ,  $\overline{LW_{IN}}$ ,  $\overline{LW_{OUT}}$ ,  $\overline{R_n}$  and  $\overline{LE}$  at daily scale, and reduced the RMSE by ~~4.62~~ 0.24 W/m<sup>2</sup> to ~~14.64~~ 0.48 W/m<sup>2</sup> and increased the R<sup>2</sup> by ~~0.04~~ 0.01 to ~~0.41~~ 0.38 for  $\overline{SW_{IN}}$ ,  $\overline{LW_{IN}}$ ,  $\overline{LW_{OUT}}$ ,  $\overline{R_n}$ ,  $\overline{LE}$  and  $\overline{H}$  at 8-day scale, when these estimates were validated against in situ observations at ~~134-44 EC independent test sites.~~ Furthermore, the CoSEB-based datasets effectively captured the spatial-temporal variability of global land surface

radiation and heat fluxes, aligning well with those from the mainstream products.

Our developed datasets hold significant potential for application across diverse fields such as agriculture, forestry, hydrology, meteorology, ecology, and environmental science. They can facilitate comprehensive studies on the variability, impacts, responses, adaptation strategies, and mitigation measures of global and regional land surface radiation and heat fluxes under the influences of climate change and human activities. These datasets will provide valuable insights and data support for scientific research, policy-making, and environmental management, advancing global solutions to address climate change.

### **Author contribution**

JW: Writing – original draft, Visualization, Software, Formal analysis, Data curation. RT: Writing – original draft, Validation, Supervision, Methodology, Funding acquisition, Formal analysis, Conceptualization. ML: Writing – review & editing, Validation. ZL: Writing – review & editing.

### **Competing interests**

The authors declare that they have no conflict of interest.

### **Acknowledgment**

We thank the work from the AmeriFlux, FLUXNET, EuroFlux, OzFlux, ChinaFLUX, the National Tibetan Plateau/Third Pole Environment Data Center and SURFRAD for providing in situ measurements. We would also like to thank Dr. Martin Jung and Dr. Ulrich Weber for providing the FLUXCOM Bowen ratio-corrected products. This work is supported by the National Natural Science Foundation of China [42271378], and the Strategic Priority Research Program of the Chinese Academy of Sciences (Grant No. XDB0740202).

---

## 817 **References**

- 818 Bai, Y., Mallick, K., Hu, T., Zhang, S., Yang, S. and Ahmadi, A.: Integrating machine  
819 learning with thermal-driven analytical energy balance model improved  
820 terrestrial evapotranspiration estimation through enhanced surface conductance,  
821 *Remote Sens. Environ.*, 311, 114308. 10.1016/j.rse.2024.114308, 2024.
- 822 Bartkowiak, P., Ventura, B., Jacob, A. and Castelli, M.: A Copernicus-based  
823 evapotranspiration dataset at 100 m spatial resolution over four Mediterranean  
824 basins, *Earth Syst. Sci. Data*, 16, 4709-4734. 10.5194/essd-16-4709-2024, 2024.
- 825 Berbery, E. H., Mitchell, K. E., Benjamin, S., Smirnova, T., Ritchie, H., Hogue, R. and  
826 Radeva, E.: Assessment of land - surface energy budgets from regional and  
827 global models, *J. Geophys. Res.-Atmos.*, 104, 19329-19348.  
828 10.1029/1999jd900128, 1999.
- 829 Betts, A. K., Ball, J. H., Beljaars, A. C. M., Miller, M. J. and Viterbo, P. A.: The land  
830 surface - atmosphere interaction: A review based on observational and global  
831 modeling perspectives, *J. Geophys. Res.-Atmos.*, 101, 7209-7225.  
832 10.1029/95jd02135, 1996.
- 833 Castelli, M., Anderson, M. C., Yang, Y., Wohlfahrt, G., Bertoldi, G., Niedrist, G.,  
834 Hammerle, A., Zhao, P., Zebisch, M. and Notarnicola, C.: Two-source energy  
835 balance modeling of evapotranspiration in Alpine grasslands, *Remote Sens.*  
836 *Environ.*, 209, 327-342. 10.1016/j.rse.2018.02.062, 2018.
- 837 Chen, J., He, T., Jiang, B. and Liang, S.: Estimation of all-sky all-wave daily net  
838 radiation at high latitudes from MODIS data, *Remote Sens. Environ.*, 245,  
839 111842. 10.1016/j.rse.2020.111842, 2020.
- 840 de Wit, A. J. W., Boogaard, H. L. and van Diepen, C. A.: Spatial resolution of  
841 precipitation and radiation: The effect on regional crop yield forecasts, *Agric.*  
842 *For. Meteorol.*, 135, 156-168. 10.1016/j.agrformet.2005.11.012, 2005.
- 843 Du, Y., Wang, T., Zhou, Y., Letu, H., Li, D. and Xian, Y.: Towards user-friendly all-sky  
844 surface longwave downward radiation from space: General scheme and product,  
845 *Bull. Amer. Meteorol. Soc.*, 105, E1303–E1319. 10.1175/bams-d-23-0126.1,  
846 2024.
- 847 ElGhawi, R., Kraft, B., Reimers, C., Reichstein, M., Körner, M., Gentine, P. and  
848 Winkler, A. J.: Hybrid modeling of evapotranspiration: inferring stomatal and  
849 aerodynamic resistances using combined physics-based and machine learning,  
850 *Environ. Res. Lett.*, 18, 034039. 10.1088/1748-9326/acbbe0, 2023.
- 851 Ersi, C., Sudu, B., Song, Z., Bao, Y., Wei, S., Zhang, J., Tong, Z., Liu, X., Le, W. and  
852 Rina, S.: The potential of NIRvP in estimating evapotranspiration, *Remote Sens.*  
853 *Environ.*, 315, 114405. 10.1016/j.rse.2024.114405, 2024.
- 854 Han, Q., Zeng, Y., Zhang, L., Wang, C., Prikaziuk, E., Niu, Z. and Su, B.: Global long  
855 term daily 1 km surface soil moisture dataset with physics informed machine  
856 learning, *Sci. Data*, 10, 101. 10.1038/s41597-023-02011-7, 2023.
- 857 Huang, J., Yu, H., Guan, X., Wang, G. and Guo, R.: Accelerated dryland expansion



858 under climate change, *Nat. Clim. Chang.*, 6, 166-171. 10.1038/nclimate2837,  
859 2015.

860 Huang, L., Luo, Y., Chen, J. M., Tang, Q., Steenhuis, T., Cheng, W. and Shi, W.:  
861 Satellite-based near-real-time global daily terrestrial evapotranspiration  
862 estimates, *Earth Syst. Sci. Data*, 16, 3993-4019. 10.5194/essd-16-3993-2024,  
863 2024.

864 Jia, B., Xie, Z., Dai, A., Shi, C. and Chen, F.: Evaluation of satellite and reanalysis  
865 products of downward surface solar radiation over East Asia: Spatial and  
866 seasonal variations, *J. Geophys. Res.-Atmos.*, 118, 3431-3446.  
867 10.1002/jgrd.50353, 2013.

868 Jiang, B., Zhang, Y., Liang, S., Wohlfahrt, G., Arain, A., Cescatti, A., Georgiadis, T., Jia,  
869 K., Kiely, G., Lund, M., Montagnani, L., Magliulo, V., Ortiz, P. S., Oechel, W.,  
870 Vaccari, F. P., Yao, Y. and Zhang, X.: Empirical estimation of daytime net  
871 radiation from shortwave radiation and ancillary information, *Agric. For.*  
872 *Meteorol.*, 211-212, 23-36. 10.1016/j.agrformet.2015.05.003, 2015.

873 Jiao, B., Su, Y., Li, Q., Manara, V. and Wild, M.: An integrated and homogenized global  
874 surface solar radiation dataset and its reconstruction based on a convolutional  
875 neural network approach, *Earth Syst. Sci. Data*, 15, 4519-4535. 10.5194/essd-  
876 15-4519-2023, 2023.

877 Jung, M., Koirala, S., Weber, U., Ichii, K., Gans, F., Camps-Valls, G., Papale, D.,  
878 Schwalm, C., Tramontana, G. and Reichstein, M.: The FLUXCOM ensemble of  
879 global land-atmosphere energy fluxes, *Sci. Data*, 6, 74. 10.1038/s41597-019-  
880 0076-8, 2019.

881 Kim, Y., Park, H., Kimball, J. S., Colliander, A. and McCabe, M. F.: Global estimates  
882 of daily evapotranspiration using SMAP surface and root-zone soil moisture,  
883 *Remote Sens. Environ.*, 298, 113803. 10.1016/j.rse.2023.113803, 2023.

884 Li, B., Ryu, Y., Jiang, C., Dechant, B., Liu, J., Yan, Y. and Li, X.: BESSv2.0: A satellite-  
885 based and coupled-process model for quantifying long-term global land-  
886 atmosphere fluxes, *Remote Sens. Environ.*, 295, 113696.  
887 10.1016/j.rse.2023.113696, 2023.

888 Liang, S., Wang, D., He, T. and Yu, Y.: Remote sensing of earth's energy budget:  
889 synthesis and review, *Int. J. Digit. Earth*, 12, 737-780.  
890 10.1080/17538947.2019.1597189, 2019.

891 Liang, S., Zheng, T., Liu, R., Fang, H., Tsay, S. C. and Running, S.: Estimation of  
892 incident photosynthetically active radiation from Moderate Resolution Imaging  
893 Spectrometer data, *J. Geophys. Res.-Atmos.*, 111. 10.1029/2005jd006730, 2006.

894 Liu, S., Xu, Z., Song, L., Zhao, Q., Ge, Y., Xu, T., Ma, Y., Zhu, Z., Jia, Z. and Zhang,  
895 F.: Upscaling evapotranspiration measurements from multi-site to the satellite  
896 pixel scale over heterogeneous land surfaces, *Agric. For. Meteorol.*, 230, 97-113.  
897 10.1016/j.agrformet.2016.04.008, 2016.

898 Mohan, M. M. P., Kanchirapuzha, R. and Varma, M. R. R.: Review of approaches for  
899 the estimation of sensible heat flux in remote sensing-based evapotranspiration

models, *J. Appl. Remote Sens.*, 14, 041501-041501. 10.1117/1.Jrs.14.041501, 2020.

Mu, Q., Zhao, M. and Running, S. W.: Improvements to a MODIS global terrestrial evapotranspiration algorithm, *Remote Sens. Environ.*, 115, 1781-1800. 10.1016/j.rse.2011.02.019, 2011.

Mueller, R. W., Matsoukas, C., Gratzki, A., Behr, H. D. and Hollmann, R.: The CM-SAF operational scheme for the satellite based retrieval of solar surface irradiance — A LUT based eigenvector hybrid approach, *Remote Sens. Environ.*, 113, 1012-1024. 10.1016/j.rse.2009.01.012, 2009.

Muñoz-Sabater, J., Dutra, E., Agustí-Panareda, A., Albergel, C., Arduini, G., Balsamo, G., Boussetta, S., Choulga, M., Harrigan, S., Hersbach, H., Martens, B., Miralles, D. G., Piles, M., Rodríguez-Fernández, N. J., Zsoter, E., Buontempo, C. and Thépaut, J.-N.: ERA5-Land: a state-of-the-art global reanalysis dataset for land applications, *Earth Syst. Sci. Data*, 13, 4349-4383. 10.5194/essd-13-4349-2021, 2021.

Nemani, R. R., Keeling, C. D., Hashimoto, H., Jolly, W. M., Piper, S. C., Tucker, C. J., Myneni, R. B. and Running, S. W.: Climate-driven increases in global terrestrial net primary production from 1982 to 1999, *Science*, 300, 1560-1563. 10.1126/science.1082750, 2003.

O., S. and Orth, R.: Global soil moisture data derived through machine learning trained with in-situ measurements, *Sci. Data*, 8. 10.1038/s41597-021-00964-1, 2021.

Peng, Z., Letu, H., Wang, T., Shi, C., Zhao, C., Tana, G., Zhao, N., Dai, T., Tang, R., Shang, H., Shi, J. and Chen, L.: Estimation of shortwave solar radiation using the artificial neural network from Himawari-8 satellite imagery over China, *Journal of Quantitative Spectroscopy and Radiative Transfer*, 240, 106672. 10.1016/j.jqsrt.2019.106672, 2020.

Rios, G. and Ramamurthy, P.: A novel model to estimate sensible heat fluxes in urban areas using satellite-derived data, *Remote Sens. Environ.*, 270, 112880. 10.1016/j.rse.2021.112880, 2022.

Ryu, Y., Jiang, C., Kobayashi, H. and Detto, M.: MODIS-derived global land products of shortwave radiation and diffuse and total photosynthetically active radiation at 5 km resolution from 2000, *Remote Sens. Environ.*, 204, 812-825. 10.1016/j.rse.2017.09.021, 2018.

Sellers, P. J., Dickinson, R. E., Randall, D. A., Betts, A. K., Hall, F. G., Berry, J. A., Collatz, G. J., Denning, A. S., Mooney, H. A., Nobre, C. A., Sato, N., Field, C. B. and Henderson-Sellers, A.: Modeling the Exchanges of Energy, Water, and Carbon Between Continents and the Atmosphere, *Science*, 275, 502-509. 10.1126/science.275.5299.502, 1997.

Sun, S., Bi, Z., Xiao, J., Liu, Y., Sun, G., Ju, W., Liu, C., Mu, M., Li, J., Zhou, Y., Li, X., Liu, Y. and Chen, H.: A global 5 km monthly potential evapotranspiration dataset (1982–2015) estimated by the Shuttleworth–Wallace model, *Earth Syst. Sci. Data*, 15, 4849-4876. 10.5194/essd-15-4849-2023, 2023.

- 942 Tang, R., Peng, Z., Liu, M., Li, Z.-L., Jiang, Y., Hu, Y., Huang, L., Wang, Y., Wang, J.,  
943 Jia, L., Zheng, C., Zhang, Y., Zhang, K., Yao, Y., Chen, X., Xiong, Y., Zeng, Z.  
944 and Fisher, J. B.: Spatial-temporal patterns of land surface evapotranspiration  
945 from global products, *Remote Sens. Environ.*, 304, 114066.  
946 10.1016/j.rse.2024.114066, 2024a.
- 947 Tang, R., Wang, J., Liu, M. and Li, Z.-L.: Energy-conservation datasets of global land  
948 surface radiation and heat fluxes from 2000-2020 generated by CoSEB,  
949 National Tibetan Plateau / Third Pole Environment Data Center. [data set],  
950 <https://doi.org/10.11888/Terre.tpd.c.302559>, 2025a.
- 951 Tang, R., Wang, J., Liu, M. and Li, Z.-L.: Energy-conservation datasets of global land  
952 surface radiation and heat fluxes from 2000-2020 generated by CoSEB, Science  
953 Data Bank: Science Data Bank. [data set],  
954 <https://doi.org/10.57760/sciencedb.27228>, 2025b.
- 955 Tang, W., He, J., Qi, J. and Yang, K.: A dense station-based, long-term and high-  
956 accuracy dataset of daily surface solar radiation in China, *Earth Syst. Sci. Data*,  
957 15, 4537-4551. 10.5194/essd-15-4537-2023, 2023.
- 958 Tang, W., He, J., Shao, C., Song, J., Yuan, Z. and Yan, B.: Constructing a long-term  
959 global dataset of direct and diffuse radiation (10 km, 3 h, 1983–2018) separating  
960 from the satellite-based estimates of global radiation, *Remote Sens. Environ.*,  
961 311, 114292. 10.1016/j.rse.2024.114292, 2024b.
- 962 Tang, W., Yang, K., Qin, J., Li, X. and Niu, X.: A 16-year dataset (2000–2015) of high-  
963 resolution (3 h, 10 km) global surface solar radiation, *Earth Syst. Sci. Data*, 11,  
964 1905-1915. 10.5194/essd-11-1905-2019, 2019.
- 965 Twine, T. E., Kustas, W. P., Norman, J. M., Cook, D. R., Houser, P. R., Meyers, T. P.,  
966 Prueger, J. H., Starks, P. J. and Wesely, M. L.: Correcting eddy-covariance flux  
967 underestimates over a grassland, *Agric. For. Meteorol.*, 103, 279-300.  
968 10.1016/S0168-1923(00)00123-4, 2000.
- 969 van der Tol, C.: Validation of remote sensing of bare soil ground heat flux, *Remote Sens.*  
970 *Environ.*, 121, 275-286. 10.1016/j.rse.2012.02.009, 2012.
- 971 Wang, D., Liang, S., He, T. and Shi, Q.: Estimation of Daily Surface Shortwave Net  
972 Radiation From the Combined MODIS Data, *IEEE Trans. Geosci. Remote*  
973 *Sensing*, 53, 5519-5529. 10.1109/tgrs.2015.2424716, 2015.
- 974 Wang, D., Liang, S., Li, R. and Jia, A.: A synergic study on estimating surface  
975 downward shortwave radiation from satellite data, *Remote Sens. Environ.*, 264,  
976 112639. 10.1016/j.rse.2021.112639, 2021.
- 977 Wang, J., Tang, R., Liu, M., Jiang, Y., Huang, L. and Li, Z.-L.: Coordinated estimates  
978 of 4-day 500 m global land surface energy balance components, *Remote Sens.*  
979 *Environ.*, 326, 114795. 10.1016/j.rse.2025.114795, 2025.
- 980 Wang, K. C., Dickinson, R. E., Wild, M. and Liang, S.: Atmospheric impacts on climatic  
981 variability of surface incident solar radiation, *Atmos. Chem. Phys.*, 12, 9581-  
982 9592. 10.5194/acp-12-9581-2012, 2012.
- 983 Wang, T., Shi, J., Ma, Y., Letu, H. and Li, X.: All-sky longwave downward radiation

984 from satellite measurements: General parameterizations based on LST, column  
985 water vapor and cloud top temperature, *ISPRS-J. Photogramm. Remote Sens.*,  
986 161, 52-60. 10.1016/j.isprsjprs.2020.01.011, 2020.

987 Wang, Y., Hu, J., Li, R., Song, B. and Hailemariam, M.: Remote sensing of daily  
988 evapotranspiration and gross primary productivity of four forest ecosystems in  
989 East Asia using satellite multi-channel passive microwave measurements, *Agric.*  
990 *For. Meteorol.*, 339, 109595. 10.1016/j.agrformet.2023.109595, 2023.

991 Wild, M.: Global dimming and brightening: A review, *J. Geophys. Res.-Atmos.*, 114,  
992 D00D16. 10.1029/2008jd011470, 2009.

993 Wild, M., Folini, D., Schär, C., Loeb, N., Dutton, E. G. and König-Langlo, G.: The  
994 global energy balance from a surface perspective, *Clim. Dyn.*, 40, 3107-3134.  
995 10.1007/s00382-012-1569-8, 2012.

996 Wild, M. and Liepert, B.: The Earth radiation balance as driver of the global  
997 hydrological cycle, *Environ. Res. Lett.*, 0, 025203. 10.1088/1748-  
998 9326/5/2/025003, 2010.

999 Xia, X. A., Wang, P. C., Chen, H. B. and Liang, F.: Analysis of downwelling surface  
1000 solar radiation in China from National Centers for Environmental Prediction  
1001 reanalysis, satellite estimates, and surface observations, *J. Geophys. Res.-*  
1002 *Atmos.*, 111, D09103. 10.1029/2005jd006405, 2006.

1003 Xian, Y., Wang, T., Leng, W., Letu, H., Shi, J., Wang, G., Yan, X. and Yuan, H.: Can  
1004 Topographic Effects on Solar Radiation Be Ignored: Evidence From the Tibetan  
1005 Plateau, *Geophys. Res. Lett.*, 51, e2024GL108653. 10.1029/2024gl108653,  
1006 2024.

1007 Xu, J., Liang, S. and Jiang, B.: A global long-term (1981–2019) daily land surface  
1008 radiation budget product from AVHRR satellite data using a residual  
1009 convolutional neural network, *Earth Syst. Sci. Data*, 14, 2315-2341.  
1010 10.5194/essd-14-2315-2022, 2022a.

1011 Xu, J., Liang, S., Ma, H. and He, T.: Generating 5 km resolution 1981–2018 daily global  
1012 land surface longwave radiation products from AVHRR shortwave and  
1013 longwave observations using densely connected convolutional neural networks,  
1014 *Remote Sens. Environ.*, 280, 113223. 10.1016/j.rse.2022.113223, 2022b.

1015 Yao, Y., Liang, S., Li, X., Chen, J., Liu, S., Jia, K., Zhang, X., Xiao, Z., Fisher, J. B.,  
1016 Mu, Q., Pan, M., Liu, M., Cheng, J., Jiang, B., Xie, X., Grünwald, T., Bernhofer,  
1017 C. and Rouspard, O.: Improving global terrestrial evapotranspiration estimation  
1018 using support vector machine by integrating three process-based algorithms,  
1019 *Agric. For. Meteorol.*, 242, 55-74. 10.1016/j.agrformet.2017.04.011, 2017.

1020 Yu, L., Qiu, G. Y., Yan, C., Zhao, W., Zou, Z., Ding, J., Qin, L. and Xiong, Y.: A global  
1021 terrestrial evapotranspiration product based on the three-temperature model  
1022 with fewer input parameters and no calibration requirement, *Earth Syst. Sci.*  
1023 *Data*, 14, 3673-3693. 10.5194/essd-14-3673-2022, 2022.

1024 Zhang, C., Long, D., Zhang, Y., Anderson, M. C., Kustas, W. P. and Yang, Y.: A decadal  
1025 (2008–2017) daily evapotranspiration data set of 1 km spatial resolution and

1026 spatial completeness across the North China Plain using TSEB and data fusion,  
 1027 Remote Sens. Environ., 262, 112519. 10.1016/j.rse.2021.112519, 2021.  
 1028 Zhang, J., Zhao, L., Deng, S., Xu, W. and Zhang, Y.: A critical review of the models  
 1029 used to estimate solar radiation, Renew. Sust. Energ. Rev., 70, 314-329.  
 1030 10.1016/j.rser.2016.11.124, 2017.  
 1031 Zhang, K., Kimball, J. S., Nemani, R. R. and Running, S. W.: A continuous satellite -  
 1032 derived global record of land surface evapotranspiration from 1983 to 2006,  
 1033 Water Resour. Res., 46, W09522. 10.1029/2009wr008800, 2010.  
 1034 Zhang, X., Liang, S., Zhou, G., Wu, H. and Zhao, X.: Generating Global Land Surface  
 1035 Satellite incident shortwave radiation and photosynthetically active radiation  
 1036 products from multiple satellite data, Remote Sens. Environ., 152, 318-332.  
 1037 10.1016/j.rse.2014.07.003, 2014.  
 1038 Zhang, Y., Kong, D., Gan, R., Chiew, F. H. S., McVicar, T. R., Zhang, Q. and Yang, Y.:  
 1039 Coupled estimation of 500 m and 8-day resolution global evapotranspiration  
 1040 and gross primary production in 2002–2017, Remote Sens. Environ., 222, 165-  
 1041 182. 10.1016/j.rse.2018.12.031, 2019.  
 1042 Zheng, C., Jia, L. and Hu, G.: Global land surface evapotranspiration monitoring by  
 1043 ETMonitor model driven by multi-source satellite earth observations, J. Hydrol.,  
 1044 613, 128444. 10.1016/j.jhydrol.2022.128444, 2022.  
 1045