

1      **Energy-conservation datasets of global land surface radiation**  
2      **and heat fluxes from 2000-2020 generated by CoSEB**

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12     **Abstract**

13     Accurately estimating global land surface radiation [including downward  
14     shortwave radiation ( $SW_{IN}$ ), downward longwave radiation ( $LW_{IN}$ ), upward shortwave  
15     radiation ( $SW_{OUT}$ ), upward longwave radiation ( $LW_{OUT}$ ) and net radiation (Rn)] and heat  
16     fluxes [including latent heat flux (LE), soil heat flux (G) and sensible heat flux (H)] is  
17     essential for quantifying the exchange of radiation, heat and water between the land and  
18     atmosphere under global climate change. This study presents the first data-driven

19     energy-conservation datasets of global land surface radiation and heat fluxes from 2000

20     to 2020, generated by our model of Coordinated estimates of land Surface Energy

21     Balance components (CoSEB). The model that integrates GLASS and MODIS remote

22     sensing data, ERA5-Land reanalysis datasets, topographic data, CO<sub>2</sub> concentration data

23     as independent variables and in situ radiation and heat flux observations at 258 eddy

24     covariance sites worldwide as dependent variables within a multivariate random forest

25     technique to effectively learn the physics of energy conservation was renewed with a

26     combination of GLASS and MODIS remote sensing data, ERA5 Land reanalysis

27     datasets, topographic data, CO<sub>2</sub> concentration data, and observations at 258 eddy

28 ~~co~~variance sites worldwide from the AmeriFlux, FLUXNET, EuroFlux, OzFlux,  
29 ChinaFLUX and TPDC. The developed CoSEB-based datasets are strikingly  
30 advantageous in that [1] they are the first ~~RS-based data-driven~~ global datasets that  
31 satisfy both surface radiation balance ( $SW_{IN} - SW_{OUT} + LW_{IN} - LW_{OUT} = Rn$ ) and heat  
32 balance ( $LE + H + G = Rn$ ) among the eight fluxes, as demonstrated by both the  
33 radiation imbalance ratio [RIR, defined as  $100\% \times (SW_{IN} - SW_{OUT} + LW_{IN} - LW_{OUT})/Rn$ ]  
34 and energy imbalance ratio [EIR, defined as  $100\% \times (Rn - G - LE - H)/Rn$ ] of 0, [2] the  
35 radiation and heat fluxes are characterized by high accuracies, where (1) the RMSEs  
36 ( $R^2$ ) for daily estimates of  $SW_{IN}$ ,  $SW_{OUT}$ ,  $LW_{IN}$ ,  $LW_{OUT}$ ,  $Rn$ ,  $LE$ ,  $H$  and  $G$  from the  
37 CoSEB-based datasets at 44 independent test sites were 28.5137.52 W/m<sup>2</sup> (0.81),  
38 10.394.20 W/m<sup>2</sup> (0.42), 14.2922.47 W/m<sup>2</sup> (0.90), 10.623.78 W/m<sup>2</sup> (0.95), 22.409.66  
39 W/m<sup>2</sup> (0.77), 24.3830.87 W/m<sup>2</sup> (0.60), 22.679.75 W/m<sup>2</sup> (0.44) and 6.775.69 W/m<sup>2</sup>  
40 (0.44), respectively, as well as for 8-day estimates were 12.81 W/m<sup>2</sup>, 7.08 W/m<sup>2</sup>, 9.22  
41 W/m<sup>2</sup>, 8.34 W/m<sup>2</sup>, 13.38 W/m<sup>2</sup>, 19.99 W/m<sup>2</sup>, 17.44 W/m<sup>2</sup> and 4.25 W/m<sup>2</sup>, respectively,  
42 (2) the CoSEB-based datasets, in comparison to the mainstream products/datasets (i.e.  
43 GLASS, BESS-Rad, BESSV2.0, FLUXCOM, MOD16A2, PML\_V2 and ETMonitor)  
44 that generally separately estimated subsets of the eight flux components, better agreed  
45 with the in situ observations. Our developed datasets hold significant potential for  
46 application across diverse fields such as agriculture, forestry, hydrology, meteorology,  
47 ecology, and environmental science, which can facilitate comprehensive studies on the  
48 variability, impacts, responses, adaptation strategies, and mitigation measures of global  
49 and regional land surface radiation and heat fluxes under the influences of climate  
50 change and human activities. The CoSEB-based datasets are open access and available  
51 through the National Tibetan Plateau Data Center (TPDC) at  
52 <https://doi.org/10.11888/Terre.tpdc.302559> (Tang et al., 2025a) and through the Science  
53 Data Bank (ScienceDB) at <https://doi.org/10.57760/sciencedb.27228> (Tang et al.,  
54 2025b).

55 **Key words:** Surface energy balance; Surface radiation balance; Shortwave/Longwave

56 radiation; Net radiation; Sensible/Latent heat flux; Evapotranspiration; CoSEB

57 **1 Introduction**

58 Land surface radiation balance and heat balance play important roles in Earth's  
59 climate system, representing the physical processes by which the surface-atmosphere  
60 absorbs and redistributes radiation and heat fluxes (Berbery et al., 1999; Betts et al.,  
61 1996; Mueller et al., 2009; Sellers et al., 1997; Xu et al., 2022a), and facilitating the  
62 exchange of water, energy, carbon, and other agents essential to climatic and ecological  
63 systems and human society (Jia et al., 2013; Wang et al., 2012; Wild, 2009; Wild et al.,  
64 2012; Xia et al., 2006). Accurately monitoring the spatial and temporal variations of  
65 global land surface radiation [including downward shortwave radiation ( $SW_{IN}$ ),  
66 downward longwave radiation ( $LW_{IN}$ ), upward shortwave radiation ( $SW_{OUT}$ ), upward  
67 longwave radiation ( $LW_{OUT}$ ) and net radiation (Rn)] and heat fluxes [including latent  
68 heat flux (LE), soil heat flux (G) and sensible heat flux (H)] is indispensable for  
69 quantifying the exchange of radiation, heat and water between the land and atmosphere  
70 under global climate change (Ersi et al., 2024; Liang et al., 2019; Rios & Ramamurthy,  
71 2022; Tang et al., 2024a; Wang et al., 2021), and for studying solar energy utilization  
72 (Tang et al., 2024b; Zhang et al., 2017), hydrological cycle (Huang et al., 2015; Wild &  
73 Liepert, 2010), ecosystem productivity (Nemani et al., 2003), agricultural management  
74 (De Wit et al., 2005) and ecological protection (Tang et al., 2023). Remote sensing (RS)  
75 technology, with its high spatial-temporal resolution and applicability over large areas,  
76 is considered to be the most effective and economical means for obtaining global land  
77 surface radiation and heat fluxes (Liu et al., 2016; Van Der Tol, 2012; Zhang et al.,  
78 2010).

79 In past decades, numerous RS-based products/datasets of global surface radiation  
80 and heat fluxes have significantly advanced, which were generally generated by  
81 physical (Li et al., 2023; Mu et al., 2011; Yu et al., 2022) or statistical methods (Jiao et  
82 al., 2023; Jung et al., 2019; Peng et al., 2020). However, two key limitations still exist  
83 in these products. Firstly, most available products provide only a single component of

84 land surface radiation or heat fluxes, e.g. ETMonitor (Zheng et al., 2022) and  
85 MOD16A2 (Mu et al., 2011) only estimating LE, leading to the failure to satisfy surface  
86 radiation balance and heat balance when the single radiation or heat flux is utilized in  
87 conjunction with products containing other radiation and heat components (Wang et al.,  
88 2025), and further posing significant uncertainties to understand the interactions and  
89 redistributions of surface radiation and energy in the Earth-atmosphere system.  
90 Secondly, a few products, e.g., FLUXCOM (Jung et al., 2019) and GLASS (Jiang et al.,  
91 2015; Zhang et al., 2014), generated datasets for multiple components of surface  
92 radiation and heat fluxes by using independent-separate estimates from the  
93 uncoordinated models, which make them difficult to abide by surface radiation and heat  
94 conservation. These energy-imbalanced and radiation-imbalanced estimates among  
95 multiple components from previous products/datasets severely limit their in-depth  
96 applications in analyzing the spatial and temporal trends, simulating the physical  
97 processes of radiation, heat and water cycles as well as revealing the attributions and  
98 mechanisms in Earth-surface system under global climate change. It was impending  
99 and-imperative to develop global datasets of land surface radiation and heat fluxes  
100 characterized by high accuracies, radiation balance as well as heat balance, to better  
101 meet the requirements in practical applications of various fields.

102 Our proposed data-driven model/framework of Coordinated estimates of land  
103 Surface Energy Balance components (CoSEB) (Wang et al., 2025), which effectively  
104 learns the underlying physical interrelations (i.e., surface energy conservation law)  
105 among multiple targeted variables, provides an unprecedented opportunity to develop  
106 global datasets of land surface radiation and heat fluxes that can not only  
107 simultaneously provide high-accuracy estimates of these components but also adhere  
108 to surface radiation- and heat-conservation laws.

109 The objectives of this study are twofold: (1) to develop high-accuracy datasets of  
110 global land surface radiation and heat fluxes, which comply with the principles of  
111 radiation balance and heat balance, using our CoSEB model renewed based on in situ

112 observations, remote sensing data and reanalysis datasets; (2) to validate the  
113 datasets/model estimates against data from in situ observations, mainstream products  
114 as well as estimates from uncoordinated random forest (RF) techniques. Section 2  
115 introduces the data resources used in this study. Section 3 briefly describes the method  
116 we used to estimate global surface radiation and heat fluxes. Section 4 presents the  
117 evaluation of the datasets/model estimates generated by our renewed CoSEB model.  
118 Section 5 discusses the superiority, potential applications and uncertainties of the  
119 developed datasets. Data availability is given in Section 6, and a summary and  
120 conclusion is provided in Section 7.

121 **2 Data**

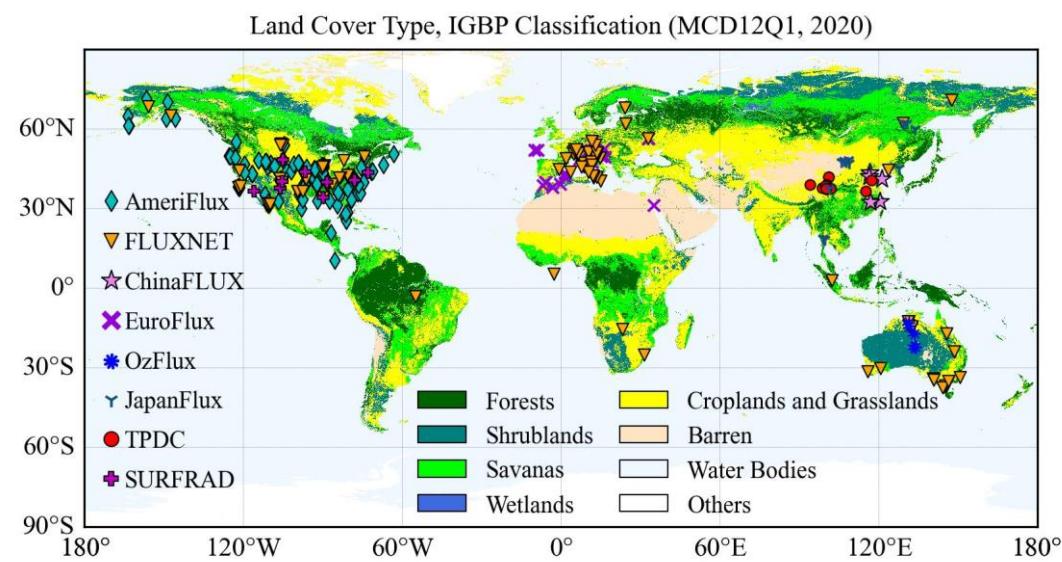
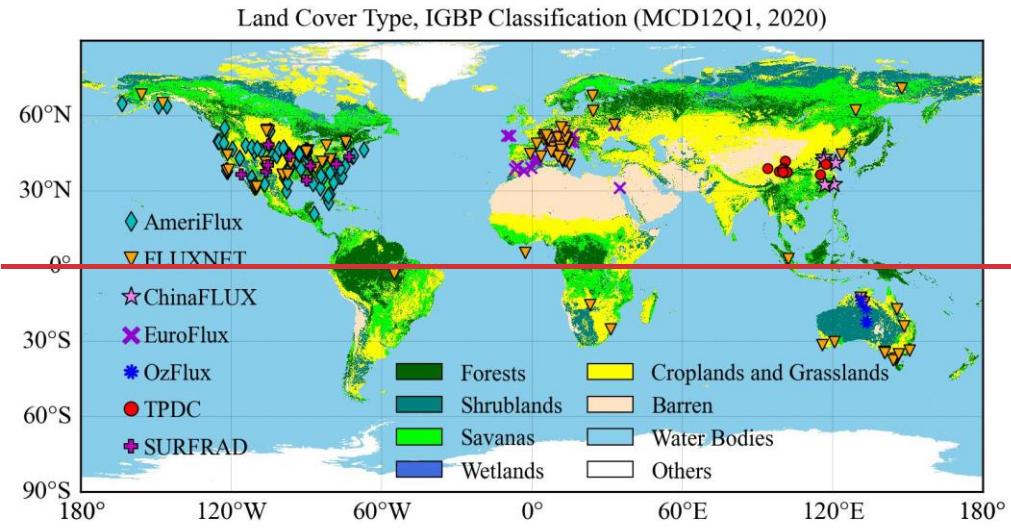
122 **2.1 Ground-based observations**

123 In this study, the in situ observations of land surface radiation and heat fluxes at  
124 258-302 eddy covariance (EC) sites from the networks of AmeriFlux (145-174 sites,  
125 2000–2020, <https://AmeriFlux.lbl.gov/Data/>, last access: 6 August 2024), EuroFlux (72  
126 sites, 2000-2020, <http://www.europe-fluxdata.eu/>, last access: 6 August 2024), OzFlux  
127 (5 sites, 2007-2012, <https://data.ozflux.org.au/>, last access: 6 August 2024), FLUXNET  
128 (108 sites, 2000–2014, <https://FLUXNET.org/Data/download-Data/>, last access: 6  
129 August 2024), JapanFlux (15 sites, 2001-2020, https://ads.nipr.ac.jp/japan-flux2024/,  
130 last access: 10 October 2025), ChinaFLUX (5 sites, 2005-2020,  
131 <http://www.chinaflux.org/>, last access: 6 August 2024) and National Tibetan  
132 Plateau/Third Pole Environment Data Center (TPDC, 13 sites, 2012–2020,  
133 <https://Data.tpdc.ac.cn/en/Data>, last access: 6 August 2024) were used (Fig. 1), where  
134 37, 48 and 5 sites in FLUXNET were also shared in AmeriFlux, EuroFlux and OzFlux,  
135 respectively. These 258-302 sites were filtered out from all collected 1008-1098 sites  
136 by following the quality-assurance and quality-control steps, including: (1) any site  
137 with a missing component of any of the  $SW_{IN}$ ,  $SW_{OUT}$ ,  $LW_{IN}$ ,  $LW_{OUT}$ , LE, H and G was  
138 excluded, reducing the 1008-1098 sites to 388-472 sites for further analysis; (2) any

139 half-hour period with missing data for any of these components was excluded; (3) the  
140 half-hourly ground-based observations with quality-control flag of 2 or 3 (bad quality)  
141 were removed but quality-control flag of 0 and 1 (good quality) were maintained; (4) a  
142 daily average of the half-hour observations was calculated for each day with greater  
143 than 80% good-quality data, further reducing the 388-472 sites to 286-355 sites; (5) the  
144 aggregated daily LE and H were corrected for energy imbalance using the Bowen ratio  
145 method when the daily energy balance closure [defined as  $(LE + H) / (Rn - G)$ ] varied  
146 between 0.2 and 1.8 following Wang et al. (2025) to exclude physically implausible  
147 measurements; (56) extreme outliers in the daily evaporative fraction were further  
148 removed by excluding values outside the 1st–99th percentile range, a common practice  
149 in flux and remote sensing studies (Bartkowiak et al., 2024; Wang et al., 2023), further  
150 reducing the 355 sites to 337 sites. outliers were discarded, corresponding to the 1 and  
151 99 quantiles of the daily evaporation fraction, further reducing the 286 sites to 268 sites.  
152 Besides, the RS data products/datasets involved in this study collocated at the sites  
153 should not be missing, finally reducing the 268-337 sites to 258-302 sites for analysis.  
154 Note that the Rn at these sites used in this study was calculated from the sum of net  
155 longwave radiation ( $LW_{IN}$  minus  $LW_{OUT}$ ) and net shortwave radiation ( $SW_{IN}$  minus  
156  $SW_{OUT}$ ), rather than using the observed Rn directly, to ensure surface radiation balance  
157 in training datasets.

158 These 258-302 sites used in this study cover a wide range of global climate regimes  
159 across 14 land cover types, including (1) evergreen needleleaf forests (ENF, 54-55 sites);  
160 (2) evergreen broadleaf forests (EBF, 11-12 sites); (3) deciduous needleleaf forests  
161 (DNF, 1-7 sites); (4) deciduous broadleaf forests (DBF, 39-40 sites); (5) mixed forests  
162 (MF, 8 sites); (6) closed shrublands (CSH, 5 sites); (7) open shrublands (OSH, 9-11  
163 sites); (8) woody savannas (WSA, 6 sites); (9) savannas (SAV, 10 sites); (10) grasslands  
164 (GRA, 54-62 sites); (11) permanent wetlands (WET, 16-22 sites); (12) croplands (CRO,  
165 43-59 sites); (13) water bodies (WAT, 1 sites); (14) cropland/natural vegetation mosaics  
166 (CVM, 1-4 sites). Among them, 44 sites (~15% of the total, see Table S1) were isolated

167 to serve as spatially independent sites to test the generated datasets and they did not  
168 participate in the development of the model/datasets.



172 Fig. 1 Spatial distribution of the 258-302 eddy covariance sites from AmeriFlux, FLUXNET,  
173 EuroFlux, OzFlux, JapanFlux, ChinaFLUX and TPDC, and nine radiation sites from  
174 SURFRAD involved for analysis in this study.

175 Furthermore, ground-based radiation observations from nine sites that are located  
176 in large flat agricultural areas covered by crops and grasses from SURFRAD were also  
177 introduced to validate land surface radiation estimates. Similar to the preprocessing  
178 performed on the observations of the 258-302 EC sites, the  $SW_{IN}$ ,  $SW_{OUT}$ ,  $LW_{IN}$ ,  $LW_{OUT}$   
179 and  $Rn$  from the SURFRAD were also quality-controlled and aggregated to daily data.

180 Spatial distribution of the 258–302 EC sites and nine radiation sites from SURFRAD  
 181 are shown in Fig. 1, with site details (latitude, longitude, land cover types, digital  
 182 elevation model and temporal coverage) provided in Supplementary Tables S1 and S2.

183 **Table 1 Summary of mainstream datasets/products for inter-comparison used in this study**

Products/ datasets	Reso- lution	Time- coverage	Variables	Algorithms	References
GLASS	0.05°/ daily	2000– 2018	$SW_{IN}$ , $LW_{IN}$ , $LW_{OUT}$ , $Rn$	Machine- learning, direct estimation algorithm	Wang et al. (2015); Xu et al. (2022b); Jiang et al. (2015)
BESS_Rad	0.05°/ daily	2000– 2020	$SW_{IN}$	BESS process- model	Ryu et al. (2018)
BESSV2.0	0.05°/ daily	2000– 2020	$Rn$ , $LE$	BESS process- model	Li et al. (2023)
FLUXCOM	0.0833°/ 8-day	2000– 2020	$Rn$ , $LE$ , $H$	Model tree- ensembles	Jung et al. (2019)
MOD16A2	500 m/ 8-day	2000– 2020	$Rn$ , $LE$	Modified Penman- Monteith equation Penman Monteith- Leuning model,	Mu et al. (2011)
PML_V2	500 m/ 8-day	2002– 2020	$LE$	Priestly Taylor- equation and Gash model Shuttleworth- Wallace two- source scheme,	Zhang et al. (2019)
ETMonitor	1 km/ daily	2000– 2020	$LE$	Gash model and Penman equation	Zheng et al. (2022)

184 **2.2 Climate/meteorology and remote sensing data**

185 To generate global datasets of land surface radiation and heat fluxes from 2000 to  
 186 2020, five types of climate/meteorology and remote sensing data were used in this study,  
 187 including:

188 (1) ERA5-Land reanalysis datasets (<https://cds.climate.copernicus.eu/>, last access: 6  
 189 August 2024) with the spatial resolution of ~9 km from 1950 (Muñoz-Sabater et  
 190 al., 2021). Following our previous work (Wang et al., 2025), this study used

variables from the ERA5-Land datasets to drive the model, including near-surface 2 m air temperature ( $T_a$ ), soil temperature in layer 1 (0-7 cm,  $T_{s1}$ ), soil volumetric moisture content in layer 1 (0-7 cm,  $SM1$ ), solar radiation reaching the surface of the earth ( $SW_{IN}^{ERA5}$ ), net thermal radiation at the surface ( $LW_{net}$ ), pressure of the atmosphere ( $PA$ ), 10 m wind speed ( $WS$ ), precipitation ( $P_r$ ) and the 2 m dewpoint temperature, daily minimum and maximum air temperature [for calculating relative air humidity ( $RH$ )].

(2) GLASS datasets (<https://glass.bnu.edu.cn/>, last access: 6 August 2024), which provide the 500 m 8-day leaf area index ( $LAI$ ) and fractional vegetation cover ( $FVC$ ) from February 2000 to December 2021.

(3) MOD44B product (<https://lpdaac.usgs.gov/>, last access: 6 August 2024), which offers yearly 250 m percent tree cover ( $PTC$ ) since 2000, representing the percentage (0~100%) of a pixel covered by tree canopy.

(4) NOAA/GML atmospheric carbon dioxide ( $CO_2$ ) concentration data, providing monthly global marine surface mean data since 1958 ([ftp://aftp.cmdl.noaa.gov/products/trends/co2/co2\\_mm\\_gl.txt](ftp://aftp.cmdl.noaa.gov/products/trends/co2/co2_mm_gl.txt), last access: 6 August 2024).-

(5) GMTED2010 topographic data ([https://topotools.cr.usgs.gov/gmted\\_viewer/gmted2010\\_global\\_grids.php](https://topotools.cr.usgs.gov/gmted_viewer/gmted2010_global_grids.php), last access: 6 August 2024), providing 500 m digital elevation model (DEM), slope, and aspect.

The ~9 km ERA5-Land datasets were spatially interpolated to 500 m using the cubic convolution method, and the 250 m PTC was resampled to 500 m using the arithmetic averaging method.

### 2.3 Mainstream datasets/products for inter-comparison

Mainstream RS-based datasets/products of moderate-resolution global land surface radiation and heat fluxes were collected for inter-comparison (Table 1),

218 including (1) the daily  $0.05^\circ$  GLASS  $SW_{IN}$ ,  $LW_{IN}$ ,  $LW_{OUT}$  and Rn products from 2000 to  
219 2018 (<https://glass.bnu.edu.cn/>, last access: 6 August 2024), (2) the daily  $0.05^\circ$   
220 Breathing Earth System Simulator Radiation (BESS-Rad)  $SW_{IN}$  products from 2000 to  
221 2020 (<https://www.environment.snu.ac.kr/bess-rad>), (3) the daily  $0.05^\circ$  BESS  
222 Version2.0 (BESSV2.0) Rn and LE products from 2000 to 2020  
223 (<https://www.environment.snu.ac.kr/bessv2>), (4) the 8-day  $0.0833^\circ$  FLUXCOM Rn, LE  
224 and H products from 2001 to 2020 (<https://fluxcom.org/>, last access: 6 August 2024),  
225 (5) the daily 1 km ETMonitor LE product from 2000 to 2020 (<https://data.casearth.cn/>,  
226 last access: 6 August 2024), (6) the 8-day 500 m Penman-Monteith-Leuning Version2  
227 (PML\_V2, <https://www.tpdc.ac.cn/>, last access: 6 August 2024) LE product from 2000  
228 to 2020; and (7) the 8-day 500 m MOD16A2 (<https://lpdaac.usgs.gov/>, last access: 6  
229 August 2024) LE product from 2000 to 2020.

230 The GLASS  $SW_{IN}$  products are derived from a combination of the GLASS  
231 broadband albedo product and the surface shortwave net radiation estimates, where the  
232 surface shortwave net radiation is estimated using linear regression with MODIS top-  
233 of-atmosphere (TOA) spectral reflectance (Wang et al., 2015). The GLASS  $LW_{IN}$  and  
234  $LW_{OUT}$  products are generated using densely connected convolutional neural networks,  
235 incorporating Advanced Very High-Resolution Radiometer (AVHRR) TOA reflectance  
236 and ERA5 near-surface meteorological data (Xu et al., 2022b). The GLASS Rn  
237 products are estimated from the meteorological variables from MERRA2 and surface  
238 variables from GLASS using the multivariate adaptive regression splines model (Jiang  
239 et al., 2015). The BESS-Rad and BESSV2.0 estimate  $SW_{IN}$  and Rn using a radiative  
240 transfer model (i.e., Forest Light Environmental Simulator, FLiES) with an artificial  
241 neural network based on MODIS and MERRA2 reanalysis datasets, and using FLiES  
242 based on MODIS products and NCEP/NCAR reanalysis data, respectively (Li et al.,  
243 2023; Ryu et al., 2018). Moreover, the BESSV2.0 (Li et al., 2023), MOD16A2 (Mu et  
244 al., 2011), PML\_V2 (Zhang et al., 2019) and ETMonitor (Zheng et al., 2022) generated  
245 global LE by physical models, such as Penman-Monteith equation, Priestley-Taylor

246 equation and/or Shuttleworth-Wallace two-source scheme. The FLUXCOM Rn, LE and  
 247 H datasets are obtained through multiple machine learning methods based on in situ  
 248 observations from FLUXNET and remote sensing and meteorological data (Jung et al.,  
 249 2019). For better consistency, RF-based 8-day 0.0833° Rn and Bowen ratio-corrected  
 250 LE and H for the periods of 2000 to 2020 from the FLUXCOM were used in this study.

251 **Table 1 Summary of mainstream datasets/products for inter-comparison used in this study**

Products/ datasets	Reso- lution	Time coverage	Variables	Algorithms	References
<u>GLASS</u>	<u>0.05°/ daily</u>	<u>2000-2018</u>	<u><math>SW_{IN}</math>, <math>LW_{IN}</math>, <math>LW_{OUT}</math>, Rn</u>	<u>Machine learning, direct estimation algorithm</u>	<u>Wang et al. (2015); Xu et al. (2022b); Jiang et al. (2015)</u>
<u>BESS-Rad</u>	<u>0.05°/ daily</u>	<u>2000-2020</u>	<u><math>SW_{IN}</math></u>	<u>BESS process model</u>	<u>Ryu et al. (2018)</u>
<u>BESSV2.0</u>	<u>0.05°/ daily</u>	<u>2000-2020</u>	<u>Rn, LE</u>	<u>BESS process model</u>	<u>Li et al. (2023)</u>
<u>FLUXCOM</u>	<u>0.0833°/ 8-day</u>	<u>2000-2020</u>	<u>Rn, LE, H</u>	<u>Model tree ensembles</u>	<u>Jung et al. (2019)</u>
<u>MOD16A2</u>	<u>500 m/ 8-day</u>	<u>2000-2020</u>	<u>LE</u>	<u>Modified Penman-Monteith equation</u>	<u>Mu et al. (2011)</u>
<u>PML_V2</u>	<u>500 m/ 8-day</u>	<u>2002-2020</u>	<u>LE</u>	<u>Penman Monteith-Leuning model, Priestly Taylor equation and Gash model</u>	<u>Zhang et al. (2019)</u>
<u>ETMonitor</u>	<u>1 km/ daily</u>	<u>2000-2020</u>	<u>LE</u>	<u>Shuttleworth-Wallace two-source scheme, Gash model and Penman equation</u>	<u>Zheng et al. (2022)</u>

252

253 **3 Methods**

254 The method used to generate global datasets of land surface radiation and heat  
 255 fluxes is based on the CoSEB model/framework, which was developed by our recently

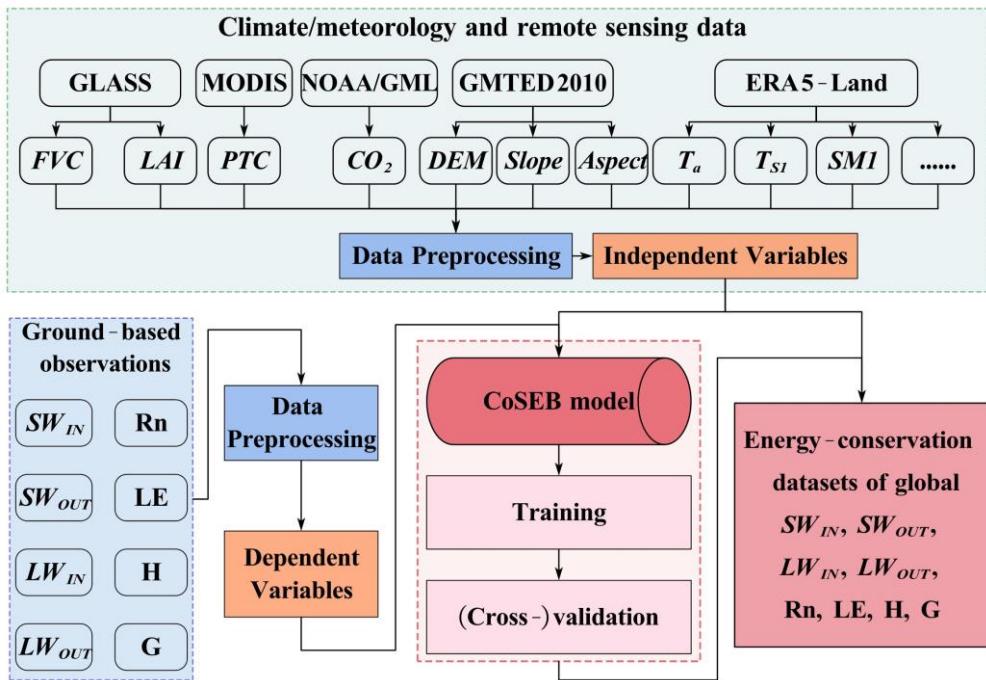
256 previously published work (Wang et al., 2025), to coordinately estimate global land  
257 surface energy balance components (including Rn, LE, H and G) using the multivariate  
258 random forest technique, with a combination of MODIS and GLASS products, ERA5-  
259 Land reanalysis datasets, and in situ observations at 336 EC sites ~~from the FLUXNET, AmeriFlux, ChinaFLUX, EuroFlux, OzFlux and Heihe River Basin flux network~~. The  
260 CoSEB model was demonstrated to be able to produce high-accuracy estimates of land  
261 surface energy components, with the RMSE of  $<17 \text{ W/m}^2$  and  $R^2$  of > 0.83 for  
262 estimating 4-day Rn, LE and H, and the RMSE of  $<5 \text{ W/m}^2$  and  $R^2$  of 0.55 for  
263 estimating 4-day G. The most praiseworthy superiority of the CoSEB model lies in its  
264 ability to balance the land surface energy components, with an energy imbalance ratio  
265 [EIR, defined as  $100\% \times (Rn - G - LE - H) / Rn$ ] of 0.

267 To coordinately estimate land surface radiation and heat fluxes that comply with  
268 both radiation balance and heat balance, one of the key procedures in the construction  
269 of the CoSEB model was to prepare training datasets that satisfy surface radiation and  
270 heat balance. For this purpose, the energy-imbalance corrections on daily in situ  
271 observed LE and H were conducted by the most widely applied Bowen ratio method  
272 [ $H^{corr} = \frac{H}{H+LE} \times (Rn - G)$ ,  $LE^{corr} = \frac{LE}{H+LE} \times (Rn - G)$ , where  $H^{corr}$  and  $LE^{corr}$   
273 represent the sensible heat flux and latent heat flux after energy-imbalance correction,  
274 respectively] with the aid of Rn and G observations, and the in situ Rn was calculated  
275 from the sum of in situ observed net longwave radiation ( $LW_{IN}$  minus  $LW_{OUT}$ ) and net  
276 shortwave radiation ( $SW_{IN}$  minus  $SW_{OUT}$ ). The input variables to renew the CoSEB  
277 model include: (1) climate/meteorology:  $T_a$ ,  $SW_{IN}^{ERA5}$ ,  $LW_{net}$ ,  $WS$ ,  $PA$ ,  $P_r$ ,  $RH$ ,  
278  $CO_2$  concentration; (2) vegetation and soil:  $LAI$ ,  $FVC$ ,  $PTC$ ,  $T_{SI} - T_{S1}$ ,  $SM1 - SM1$ ; (3)  
279 topography data: *DEM*, *Slope* and *Aspect*, in addition to longitude (*Lon*), latitude (*Lat*),  
280 and inverse relative distance from the Earth to the Sun (*dr*), in which the *dr* was  
281 calculated as  $dr = 1 + 0.033 \times \cos\left(\frac{2\pi \times DOY}{365}\right)$ , where *DOY* represents the day of year.

282 Considering that the footprint of the site-based measurements of turbulent heat fluxes  
283 is generally at a scale of hundreds of meters, to reduce the effect of differences of spatial  
284 scales between ground-based measurements (dependent variables) and remotely  
285 sensed/reanalysis datasets (independent variables), we renewed the CoSEB model at a  
286 spatial scale of 500 m for coordinately estimating global daily land surface radiation  
287 and heat fluxes, which can be expressed as follows:

288 
$$\begin{pmatrix} SW_{IN}, SW_{OUT}, LW_{IN}, \\ LW_{OUT}, Rn, LE, H, G \end{pmatrix} = f \begin{pmatrix} Lon, Lat, T_a, T_{s1}, SM1, SW_{IN}^{ERA45}, LW_{net}, PA, WS, P_r, dr \\ RH, LAI, FVC, PTC, DEM, Slope, Aspect, CO_2 \end{pmatrix} \quad (1)$$

289 To enhance model generalization, the renewed CoSEB model was reoptimized  
290 using random and grid search methods, resulting in different hyperparameters of 281  
291 decision trees, a maximum depth of 21, and minimum samples split and leaf of 8 from  
292 those of Wang et al. (2025). Site-based 10-fold cross-validation was employed to  
293 evaluate the transferability and generalization of the CoSEB model by randomly  
294 dividing all sites into ten folds, where the samples from each fold of sites in turn served  
295 as validation datasets while the remaining folds were used as training datasets, ensuring  
296 that the validation was conducted on sites spatially independent from the training data.  
297 For comparison, eight RF-based uncoordinated models for separate estimates of  $SW_{IN}$   
298  $SW_{OUT}$ ,  $LW_{IN}$ ,  $LW_{OUT}$ ,  $Rn$ ,  $LE$ ,  $H$  and  $G$  were also constructed using the same inputs as  
299 those in the renewed CoSEB model. Site-based 10-fold cross-validation was employed  
300 to assess the transferability and generalization of the CoSEB model by randomly  
301 dividing all sites into ten folds, where each fold in turn serves as validation datasets  
302 while the other folds as the training datasets, ensuring the validation of the estimates of  
303 the CoSEB was conducted at sites that are spatially independent from those selected for  
304 the training datasets. Furthermore, to benchmark the coordinated estimates from the  
305 renewed CoSEB model, eight RF-based uncoordinated models were constructed, each  
306 separately estimating one of  $SW_{IN}$ ,  $SW_{OUT}$ ,  $LW_{IN}$ ,  $LW_{OUT}$ ,  $Rn$ ,  $LE$ ,  $H$  or  $G$  using the same  
307 inputs as those in the renewed CoSEB model. Fig. 2 illustrates the flowchart for  
308 generating global datasets of land surface radiation and heat fluxes by the CoSEB model.



309

310 **Fig. 2 Flowchart for generating energy-conservation datasets of global land surface radiation**  
311 [including downward shortwave radiation ( $SW_{IN}$ ), downward longwave radiation ( $LW_{IN}$ ),  
312 upward shortwave radiation ( $SW_{OUT}$ ), upward longwave radiation ( $LW_{OUT}$ ) and net radiation  
313 ( $Rn$ ) and heat fluxes [including latent heat flux (LE), soil heat flux (G) and sensible heat flux  
314 (H)] by the CoSEB model renewed from in situ observations at 258 sites worldwide and  
315 collocated remote sensing and reanalysis datasets.

316 **4 Results**317 **4.1 Validation of the CoSEB model**318 **4.1.1 Site-based 10-fold cross-validations at 258 EC sites**

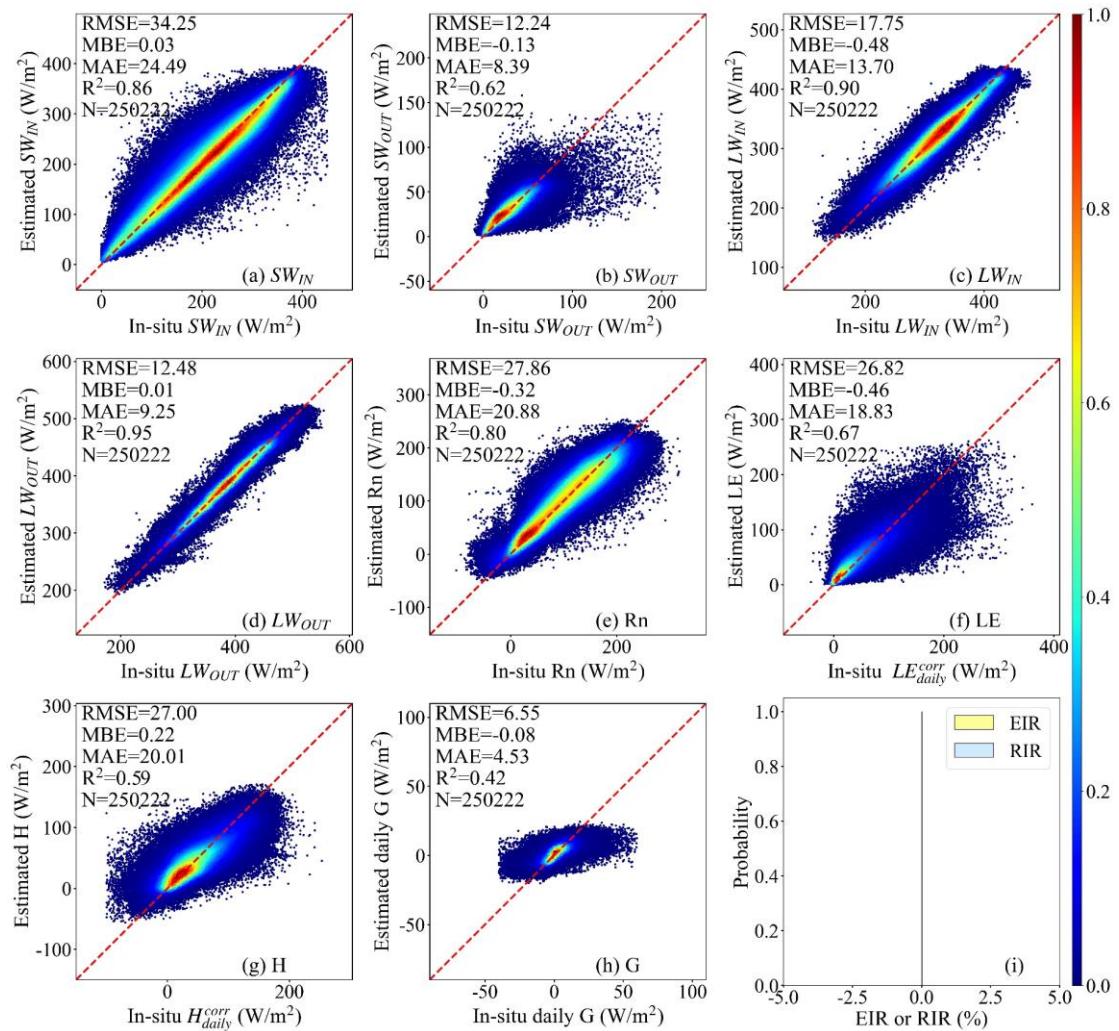
319 Fig. 3 and Fig. 4 present the scatter density plots of the site-based 10-fold cross-  
320 validation of daily  $SW_{IN}$ ,  $LW_{IN}$ ,  $SW_{OUT}$ ,  $LW_{OUT}$ ,  $Rn$ ,  $LE$ ,  $H$  and  $G$  estimated from the  
321 renewed CoSEB model and the RF-based uncoordinated models, respectively, by using  
322 the validation datasets collected at 258 EC sites worldwide. Results indicated that the  
323 estimates from both the CoSEB model and the RF-based uncoordinated models agreed  
324 well with the in situ observations, with the coefficient of determination ( $R^2$ ) varying  
325 between 0.80 and 0.95 for  $SW_{IN}$ ,  $LW_{IN}$ ,  $LW_{OUT}$  and  $Rn$ , and between 0.59 and 0.67 for  
326  $SW_{OUT}$ ,  $LE$  and  $H$ . The CoSEB model, with the root mean square error (RMSE) of 26.82  
327 to 34.25  $W/m^2$  and mean absolute error (MAE) of 18.83 to 24.49  $W/m^2$  for  $SW_{IN}$ ,  $Rn$ ,  
328  $LE$  and  $H$ , the RMSE of 12.24 to 17.75  $W/m^2$  and the MAE of 8.39 to 13.70  $W/m^2$  for

329  $SW_{OUT}$ ,  $LW_{IN}$  and  $LW_{OUT}$ , demonstrated comparable accuracies to the RF-based models,  
330 with the RMSE of 27.07 to 33.34 W/m<sup>2</sup> and MAE of 19.29 to 23.64 W/m<sup>2</sup> for  $SW_{IN}$ ,  
331 Rn, LE and H, the RMSE of 12.12 to 16.93 W/m<sup>2</sup> and the MAE of 8.68 to 12.99 W/m<sup>2</sup>  
332 for  $SW_{OUT}$ ,  $LW_{IN}$  and  $LW_{OUT}$ . In the validation of daily G, both the CoSEB and RF-based  
333 models yielded RMSEs below 7 W/m<sup>2</sup>. Comparisons with the corresponding training  
334 results (Table S3 in the Supplementary Material) indicated that although the CoSEB  
335 model performed better on the training datasets, its overall performance remained stable,  
336 suggesting that the CoSEB model was not affected by overfitting.

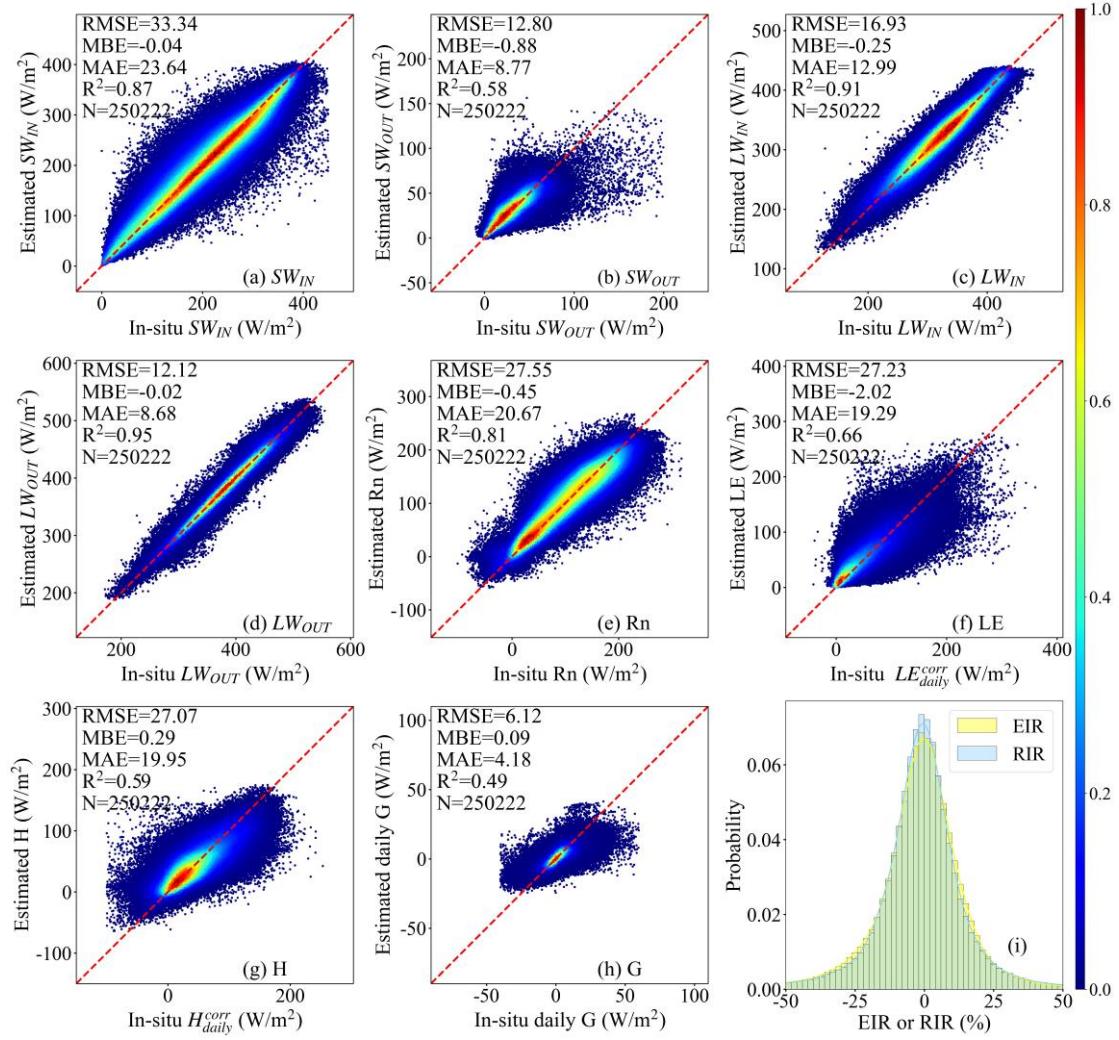
337 Strikingly, the CoSEB model exhibited large superiority in balancing the surface  
338 radiation and heat fluxes, with the radiation imbalance ratio [RIR, defined as  $100\% \times$   
339  $(SW_{IN} - SW_{OUT} + LW_{IN} - LW_{OUT})/Rn$ ]  
340  $100\% \times (SW_{IN} + LW_{IN} - SW_{OUT} - LW_{OUT} - Rn)/Rn$ ] and energy imbalance ratio [EIR,  
341 defined as  $100\% \times (Rn - G - LE - H)/Rn$   $- 100\% \times (Rn - G - LE - H)/Rn$ ] of 0, while  
342 the RF-based uncoordinated models showed substantial imbalances of the surface  
343 radiation and heat fluxes, with RIR and EIR that were approximately normally  
344 distributed, having absolute mean values of 38.84% and 31.22%, respectively, and  
345 reaching as high as 50% in some cases. Furthermore, the RIR as well as EIR tended to  
346 be higher under lower solar radiation, air temperature, or FVC, with more frequent low  
347 values of these three variables leading to a broader and less peaked distribution of RIR  
348 and EIR (see Fig. S1 in the Supplementary Material).

349 It should be pointed out that the performances of both the renewed CoSEB model  
350 and the RF based models could be further improved if the site-based 10 fold cross-  
351 validation was replaced with the sample based 10 fold cross validation (Figs. S1 and  
352 S2 in the Supplementary Material). Specifically, for the CoSEB model, using the  
353 sample based 10 fold cross validation decreased the RMSE by 0.61 to 3.92 W/m<sup>2</sup> for  
354 five radiation components and G, and by 6.25 W/m<sup>2</sup> and 5.50 W/m<sup>2</sup> for LE and H,  
355 respectively, in comparison to using the site-based 10 fold cross validation. Likewise,  
356 for the RF based models, the RMSE decreased by 1.41 to 5.25 W/m<sup>2</sup> for five radiation

357 components and G, and by  $9.63 \text{ W/m}^2$  and  $7.43 \text{ W/m}^2$  for LE and H, respectively. The  
 358  $R^2$  of both the CoSEB model and the RF based models using the sample-based 10-fold  
 359 cross-validation increased by 0.02 to 0.28 compared to the  $R^2$  using the site-based 10-  
 360 fold cross-validation.



361  
 362 **Fig. 3 Scatter density plots of the site-based 10-fold cross-validation of daily downward**  
 363 **shortwave and longwave radiation ( $SW_{IN}$  and  $LW_{IN}$ ), upward shortwave and longwave**  
 364 **radiation ( $SW_{OUT}$  and  $LW_{OUT}$ ), net radiation (Rn), soil heat flux (G), latent heat flux (LE) and**  
 365 **sensible heat flux (H) derived by the CoSEB model against in situ observed  $SW_{IN}$ ,  $LW_{IN}$ ,  $SW_{OUT}$ ,**  
 366  **$LW_{OUT}$ , Rn, G, and energy imbalance-corrected LE ( $LE_{daily}^{corr}$ ) and H ( $H_{daily}^{corr}$ ). The EIR and RIR**  
 367 **in the subfigure (i) represent the energy imbalance ratio and radiation imbalance ratio, which**  
 368 **are defined as  $100\% \times (Rn - G - LE - H)/Rn$  and  $100\% \times (Rn - G - LE - H)/Rn$  and  $100\% \times$**   
 369  **$(SW_{IN} - SW_{OUT} + LW_{IN} - LW_{OUT})/Rn$**  and  $100\% \times (SW_{IN} + LW_{IN} - SW_{OUT} - LW_{OUT} - Rn)/Rn$ ,  
 370 **respectively. The colorbar represents the normalized density of data points.**



371

372

**Fig. 4** Same as Fig. 3, but for estimates from RF-based uncoordinated models.

373

#### 4.1.2 Validation at nine radiation sites from SURFRAD

374

To further illustrate the generality and transferability of the renewed CoSEB model, the validation of estimates of the five radiation components (including  $SW_{IN}$ ,  $SW_{OUT}$ ,  $LW_{IN}$ ,  $LW_{OUT}$ ,  $Rn$ ) derived from both the CoSEB model and RF-based uncoordinated models against observations at nine radiation sites from SURFRAD was performed, as shown in Fig. 5. The results showed that both the CoSEB model and the RF-based models achieved high accuracy in estimating daily  $SW_{IN}$ ,  $SW_{OUT}$ ,  $LW_{IN}$ ,  $LW_{OUT}$  and  $Rn$ , with the RMSE of  $\sim 30$  W/m<sup>2</sup> for  $SW_{IN}$ ,  $\sim 14$  W/m<sup>2</sup> for  $SW_{OUT}$  and  $LW_{IN}$ ,  $\sim 12$  W/m<sup>2</sup> for  $LW_{OUT}$  and  $\sim 24$  W/m<sup>2</sup> for  $Rn$ , with the  $R^2 > 0.9$  for  $SW_{IN}$ ,  $LW_{IN}$  and  $LW_{OUT}$ ,  $\sim 0.65$  for  $SW_{OUT}$  and  $\sim 0.85$  for  $Rn$ . Compared to the results of the site-based 10-fold cross-validation at 258 EC sites, the performances at nine radiation sites showed slight

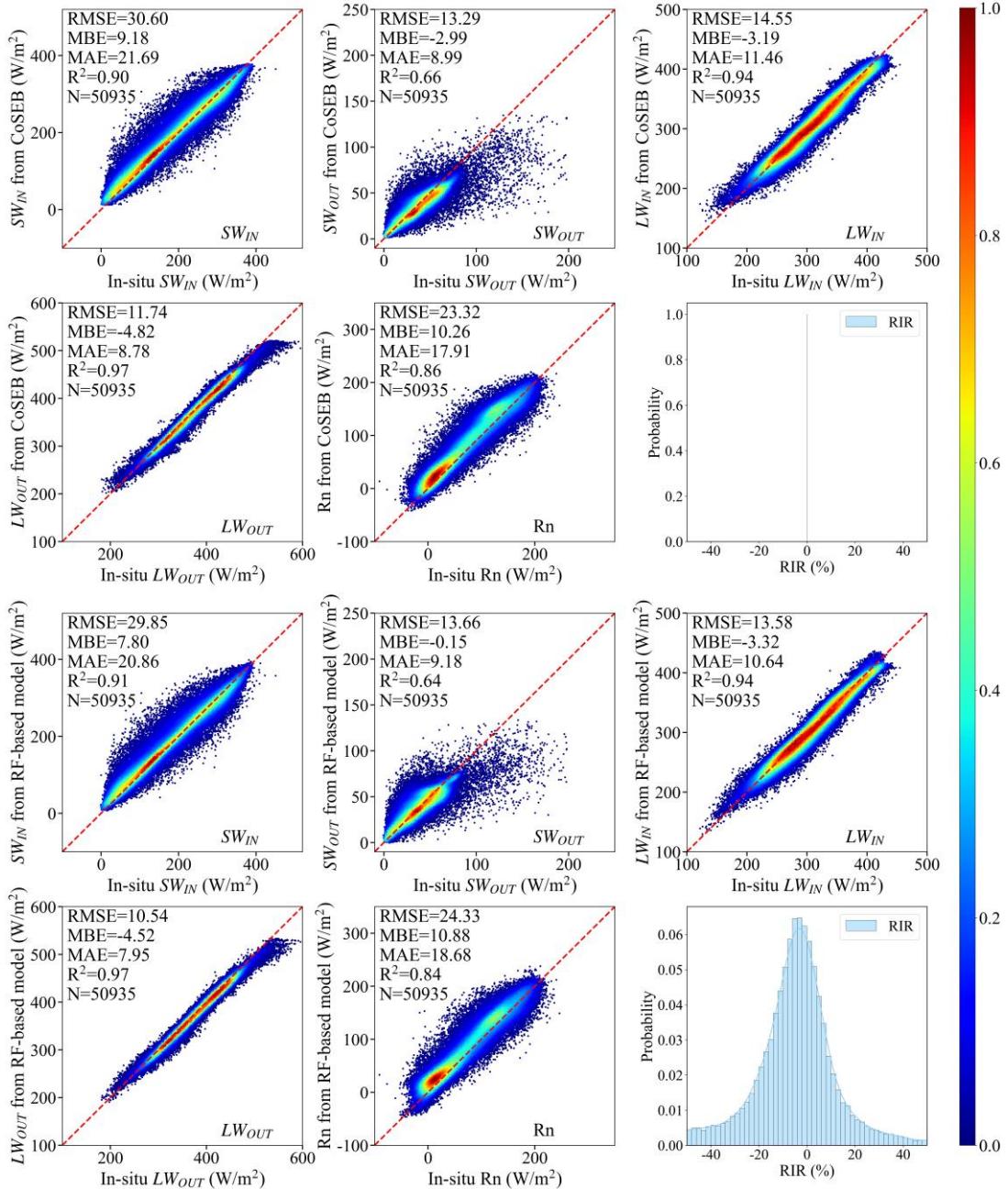
384 improvements, with the RMSE decreasing by 0.74 to 4.54 W/m<sup>2</sup> for  $SW_{IN}$ ,  $LW_{IN}$ ,  $LW_{OUT}$   
385 and  $Rn$  in the CoSEB model, but a slight degradation with the RMSE increasing by  
386  $\sim 1.05$  W/m<sup>2</sup> for  $SW_{OUT}$ , suggesting the robust performance of the CoSEB model.  
387 Furthermore, the CoSEB model demonstrated a large superiority in maintaining surface  
388 radiation balance among the five radiation components, with the RIR of 0, in contrast  
389 to the RF-based models, which failed to meet this balance, exhibiting significant RIR  
390 exceeding 50%.

#### 391 **4.2 Validation and inter-comparisons of the CoSEB-based datasets**

392 As demonstrated in Section 4.1, the renewed CoSEB model with a spatial scale of  
393 500 m achieved comparable accuracies to the RF-based uncoordinated models but  
394 outperformed them in balancing surface radiation and heat fluxes. Evidenced by the  
395 validation for its superiority, the renewed CoSEB model was then applied to the  
396 spatially aggregated input datasets to generate our developed global daily datasets with  
397 a spatial resolution of 0.05°. To further assess the performance of the developed  
398 CoSEB-based datasets, in situ observations from another 44 spatially independent test  
399 sites (see Section 2.1), which were not involved in model construction and datasets  
400 generation, were used for validation. Mainstream products (i.e. GLASS, BESS-Rad,  
401 BESSV2.0, FLUXCOM, PML\_V2, MOD16A2 and ETMonitor) were also involved for  
402 inter-comparison at the 44 test sites.

403 Note that due to the lack of moderate-resolution global RS-based products/datasets  
404 of daily and/or 8-day  $SW_{OUT}$ , H and G, the intercomparison between different  
405 products/datasets was impossible. Instead, we conducted a validation of these  
406 components from the CoSEB-based datasets against in situ observations at 44 test sites,  
407 as shown in Figs S2 and S3 in the Supplementary Material. Results indicated that the  
408 CoSEB-based datasets could provide good estimates of  $SW_{OUT}$ , H and G, with the  
409 RMSEs ( $R^2$ ) of 14.20 W/m<sup>2</sup> (0.42), 29.75 W/m<sup>2</sup> (0.44) and 5.69 W/m<sup>2</sup> (0.44) at daily  
410 scale, respectively, and the RMSE ( $R^2$ ) of 12.19 W/m<sup>2</sup> (0.39) and 4.60 W/m<sup>2</sup> (0.47) for  
411 8-day  $SW_{OUT}$  and G, respectively.





413  
 414 **Fig. 5 Scatter density plots of the validation of daily downward shortwave and longwave**  
 415 **radiation ( $SW_{IN}$  and  $LW_{IN}$ ), upward shortwave and longwave radiation ( $SW_{OUT}$  and  $LW_{OUT}$ )**  
 416 **and net radiation (Rn) from the renewed CoSEB model (upper two rows) and RF-based**  
 417 **uncoordinated models (lower two rows) -based datasets- against in situ observations at nine**  
 418 **radiation sites from SURFRAD. The RIR represents the radiation imbalance ratio, defined as**  
 419  **$100\% \times (SW_{IN} - SW_{OUT} + LW_{IN} - LW_{OUT})/Rn - 100\% \times (SW_{IN} + LW_{IN} - SW_{OUT} - LW_{OUT})/Rn$ .**

420 The colorbar represents the normalized density of data points.

421 **4.2 Validation and inter-comparisons of the CoSEB-based datasets**

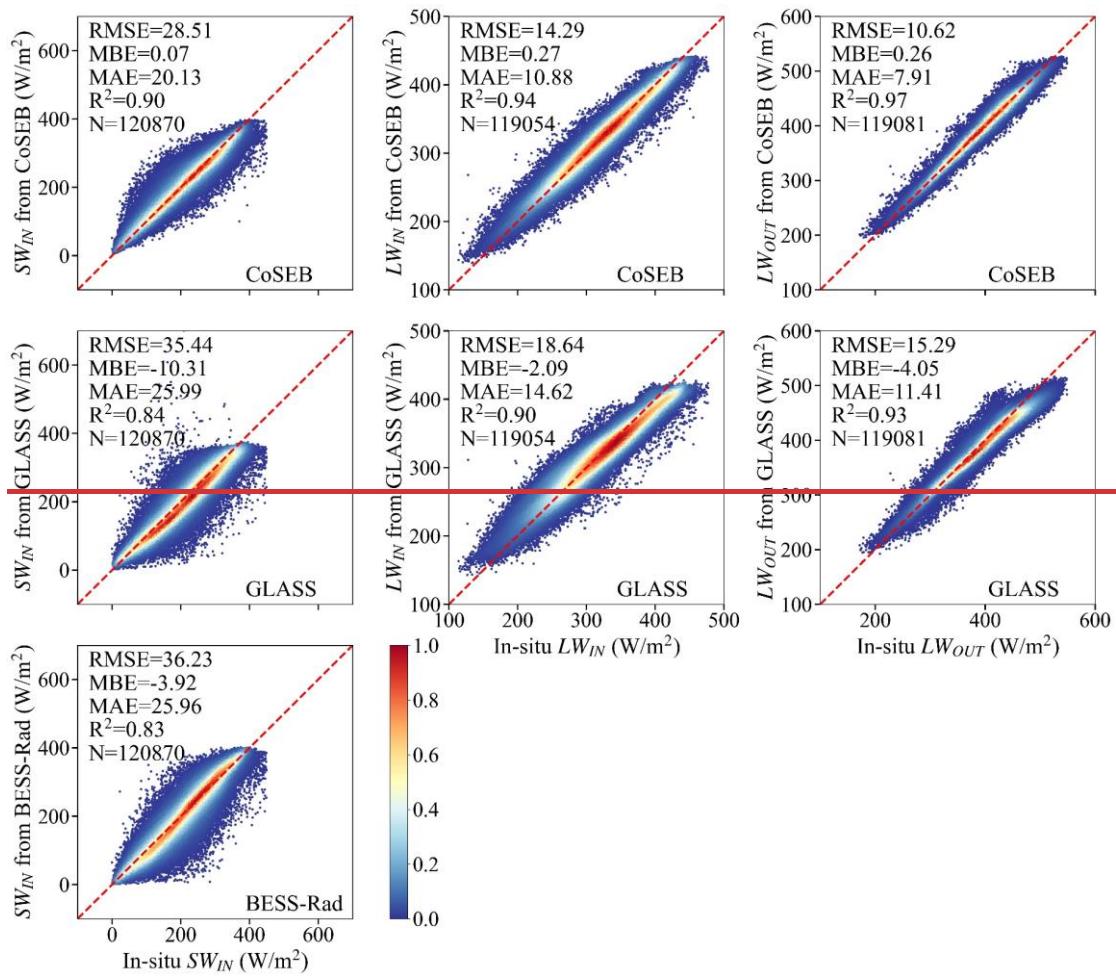
422 As demonstrated in Section 4.1, the renewed CoSEB model with a spatial scale of

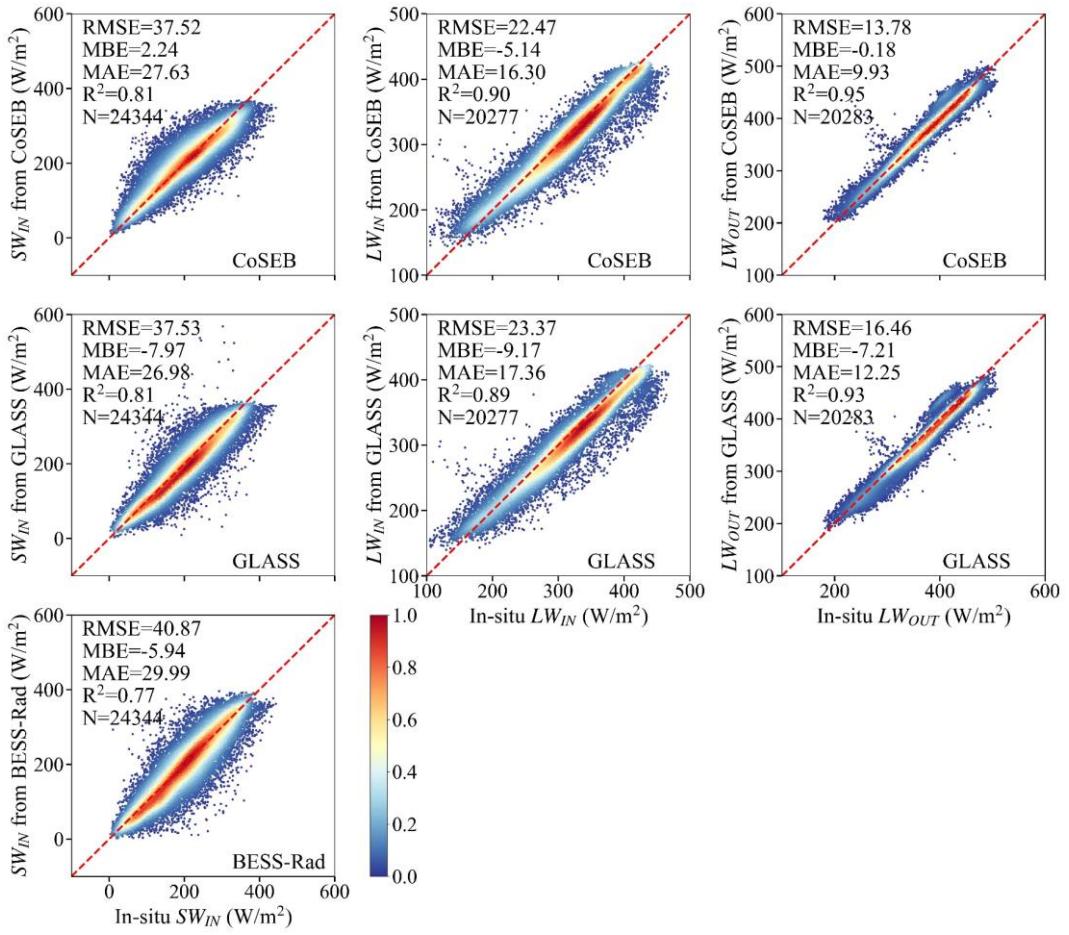
423 500 m achieved comparable accuracies to the RF-based uncoordinated models but  
424 outperformed them in balancing surface radiation and heat fluxes. Evidenced by the  
425 validation for its superiority, the renewed CoSEB model was then applied to the  
426 spatially aggregated input datasets to generate our developed global daily datasets with  
427 a spatial resolution of 0.05°. To further assess the performance of the developed datasets,  
428 in situ observations at 134 sites out of the 258 EC sites were further used to test the  
429 performance of the CoSEB-based datasets, where the 134 sites were selected based on  
430 the commonly applied criterion (Salazar Martinez et al., 2022; Tang et al., 2024a) that  
431 the fraction of the dominant land cover types (from the 500 m MCD12Q1 product)  
432 exceeded 80% within the 0.05° grid, ensuring surface homogeneity and spatial  
433 representativeness of the observations. Mainstream products (i.e. GLASS, BESS-Rad,  
434 BESSV2.0, FLUXCOM, PML\_V2, MOD16A2 and ETMonitor) were also involved for  
435 inter-comparison at the 134 EC sites.

436 Note that due to the lack of moderate resolution global RS-based products/datasets  
437 of daily and/or 8-day  $SW_{OUT}$ , H and G, the intercomparison between different  
438 products/datasets was impossible. Instead, we conducted a validation of these  
439 components from the CoSEB-based datasets against in situ observations at 134 EC sites,  
440 as shown in Figs S3 and S4 in the Supplementary Material. Results indicated that the  
441 CoSEB-based datasets could provide good estimates of  $SW_{OUT}$ , H and G, with the  
442 RMSE of 10.39 W/m<sup>2</sup>, 22.67 W/m<sup>2</sup> and 6.77 W/m<sup>2</sup> at daily scale, respectively, and the  
443 RMSE of 7.08 W/m<sup>2</sup> and 4.25 W/m<sup>2</sup> for 8-day  $SW_{OUT}$  and G, respectively.

444 Fig. 6 and Fig. 7 present the comparison of daily  $SW_{IN}$ ,  $LW_{IN}$  and  $LW_{OUT}$ , as well  
445 as Rn and LE from the CoSEB-based datasets and mainstream products/datasets  
446 (including GLASS, BESS-Rad, BESSV2.0 and ETMonitor), with in situ observations  
447 at 134 EC test sites, respectively. Overall, the estimates from the CoSEB-based  
448 datasets exhibited a closer agreement with in situ observations than those from  
449 mainstream products/datasets, where the CoSEB-based datasets reduced the RMSE by  
450 4.350.01 W/m<sup>2</sup> to 11.464.58 W/m<sup>2</sup> and increased the R<sup>2</sup> by 0.0401 to 0.309 compared

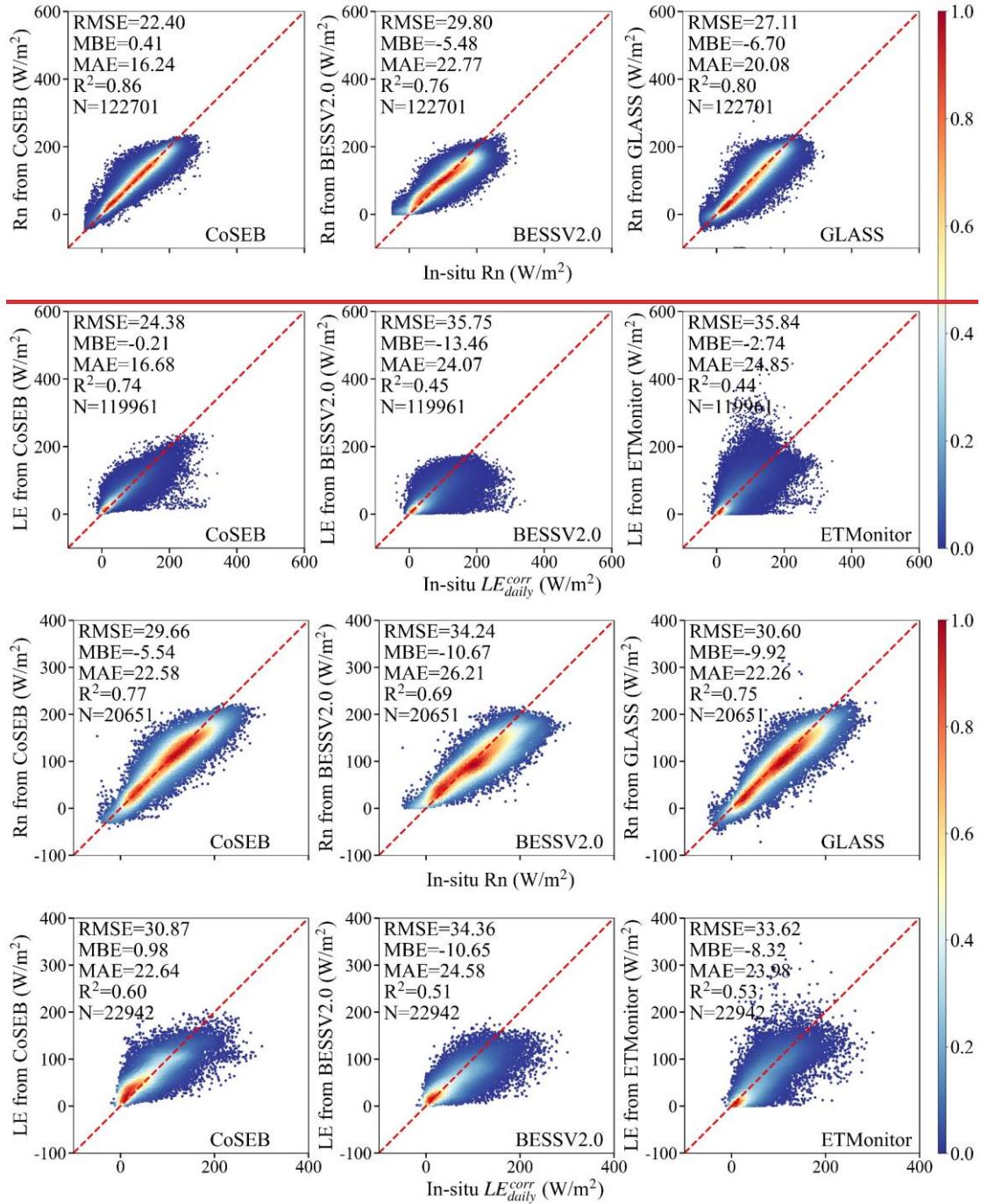
451 to mainstream products. Specifically, the RMSE for the  $SW_{IN}$ ,  $LW_{IN}$ ,  $LW_{OUT}$  increased  
 452 from 28.5137.52 W/m<sup>2</sup>, 14.2922.47 W/m<sup>2</sup> and 10.6213.78 W/m<sup>2</sup> in the CoSEB-based  
 453 datasets to 35.447.53 W/m<sup>2</sup>, 18.6423.37 W/m<sup>2</sup> and 15.2916.46 W/m<sup>2</sup> in the GLASS,  
 454 respectively, and for  $SW_{IN}$  from 28.5137.52 W/m<sup>2</sup> in the CoSEB-based datasets to  
 455 36.2340.87 W/m<sup>2</sup> in the BESS-Rad. Likewise, the RMSEs for daily Rn and LE were  
 456 22.409.66 W/m<sup>2</sup> and 24.3830.87 W/m<sup>2</sup> in the CoSEB-based datasets, which were lower  
 457 than those of 29.8034.24 W/m<sup>2</sup> and 35.754.36 W/m<sup>2</sup> in BESSV2.0, respectively, as well  
 458 as those of 27.4430.60 W/m<sup>2</sup> for Rn in GLASS and 35.843.62 W/m<sup>2</sup> for LE in  
 459 ETMonitor.





461

462 **Fig. 6 Comparison of the daily downward shortwave radiation ( $SW_{IN}$ , the first column),**  
463 **downward longwave radiation ( $LW_{IN}$ , the second column) and upward longwave radiation**  
464 **( $LW_{OUT}$ , the third column) from the CoSEB-based datasets, GLASS and BESS-Rad with the**  
465 **in situ observed  $SW_{IN}$ ,  $LW_{IN}$  and  $LW_{OUT}$  at [134-44 eddy covariance test](#) sites. The colorbar**  
466 **represents the normalized density of data points.**



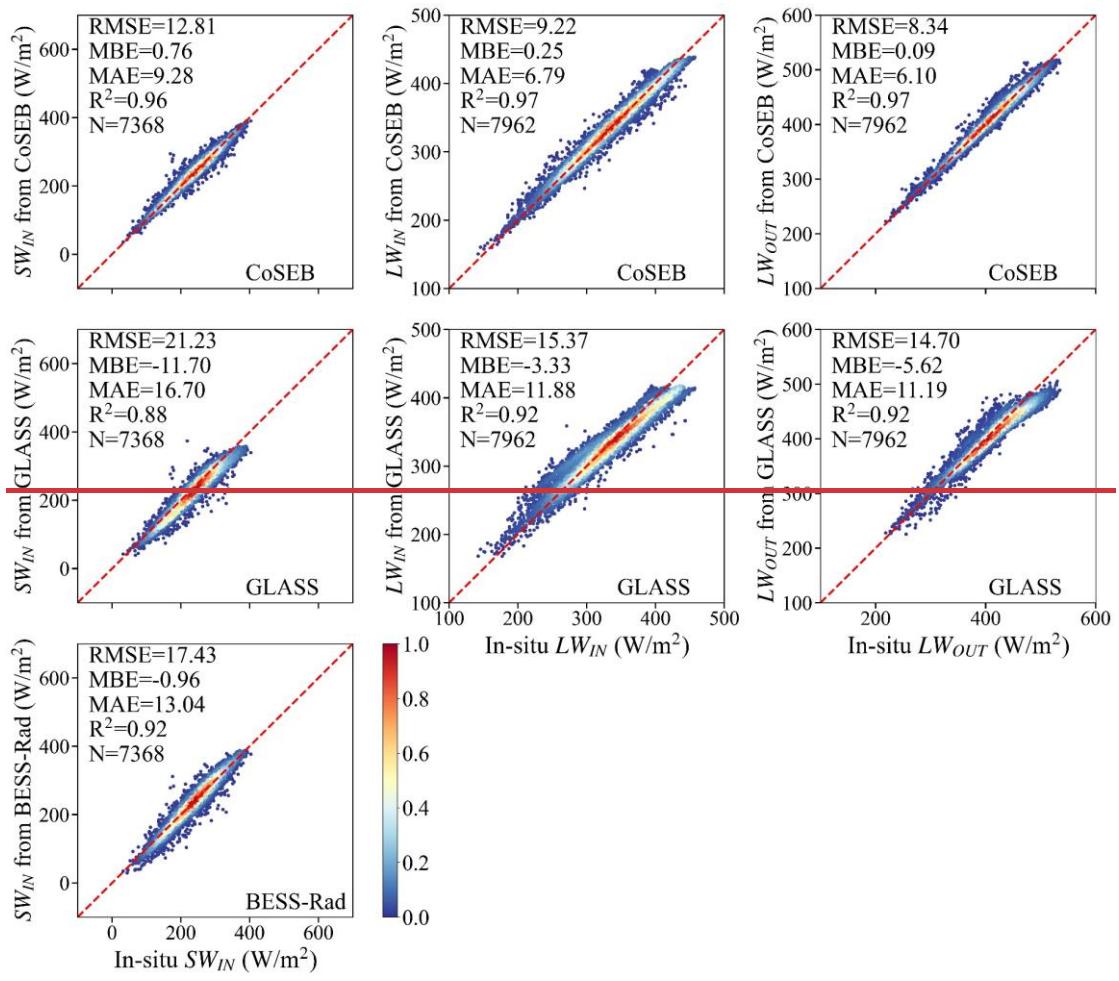
469 **Fig. 7 Comparison of the daily net radiation (Rn, the upper row) and latent heat flux (LE, the**  
 470 **lower row) from the CoSEB-based datasets, BESSV2.0, GLASS and ETMonitor with the in**  
 471 **situ observed Rn, and energy imbalance-corrected LE (LE<sup>corr</sup><sub>daily</sub>) at [134\\_44 eddy covariance test](#)**  
 472 **sites. The colorbar represents the normalized density of data points.**

473 Figs. 8, 9 and 10 compare the 8-day  $SW_{IN}$ ,  $LW_{IN}$  and  $LW_{OUT}$ , Rn and LE, as well as  
 474 H from the CoSEB-based datasets and mainstream products, with in situ observations  
 475 at [44 test134 EC](#) sites, respectively. Overall, the CoSEB-based datasets outperformed

476 the mainstream products/datasets for all surface radiation and heat fluxes, where the  
477 CoSEB-based datasets reduced the RMSE by 4.620.24 W/m<sup>2</sup> to 14.640.48 W/m<sup>2</sup> and  
478 increased the R<sup>2</sup> by 0.0401 to 0.4138 compared to mainstream products. Specifically,  
479 for  $SW_{IN}$ ,  $LW_{IN}$  and  $LW_{OUT}$ , the RMSE increased from 12.818.54 W/m<sup>2</sup>, 9.2218.50 W/m<sup>2</sup>  
480 and 8.349.41 W/m<sup>2</sup> in the CoSEB-based datasets to 21.2335 W/m<sup>2</sup>, 15.3720.39 W/m<sup>2</sup>  
481 and 14.7048 W/m<sup>2</sup> in the GLASS, respectively, and for  $SW_{IN}$  from 12.8118.54 W/m<sup>2</sup>  
482 in the CoSEB-based datasets to 17.4318.78 W/m<sup>2</sup> in the BESS-Rad. For Rn, the RMSE  
483 increased from 13.389.12 W/m<sup>2</sup> in the CoSEB-based datasets to ~23 W/m<sup>2</sup> in the  
484 FLUXCOM and GLASS and to >27 W/m<sup>2</sup> in the BESSV2.0 18.64 W/m<sup>2</sup> in the GLASS  
485 and to >23 W/m<sup>2</sup> in the FLUXCOM and BESSV2.0, while the R<sup>2</sup> decreased from 0.91  
486 82 in the CoSEB-based datasets to 0.75 in the FLUXCOM and GLASS and to 0.8262  
487 in the GLASS BESSV2.0 and to <0.72 in the FLUXCOM and BESSV2.0. Likewise, for  
488 LE, the RMSE increased from 19.9922.31 W/m<sup>2</sup> in the CoSEB-based datasets to  
489 ~26.1625 W/m<sup>2</sup> in the FLUXCOM, PML\_V2, BESSV2.0 and ETMonitor, and  
490 to >28.1732 W/m<sup>2</sup> in BESSV2.0, MOD16A2, PML\_V2 and ETMonitor, while the R<sup>2</sup>  
491 decreased from 0.867 in the CoSEB-based datasets to ~0.6560 in the FLUXCOM,  
492 PML\_V2, BESSV2.0 and ETMonitor FLUXCOM, and to <0.63 in the remaining  
493 products MOD16A1. For H, the RMSE increased from 17.4421.63 W/m<sup>2</sup> in the CoSEB-  
494 based datasets to 23.962.64 W/m<sup>2</sup> in the FLUXCOM.

495 The differences between the estimates from the CoSEB-based datasets and  
496 mainstream datasets are likely multifactorial, arising from the simplification and  
497 parameterization uncertainties in physics-based models, as well as the lack of physical  
498 constraints, limited training samples, and incomplete consideration of influencing  
499 factors in other machine-learning-based models.

500



502

503

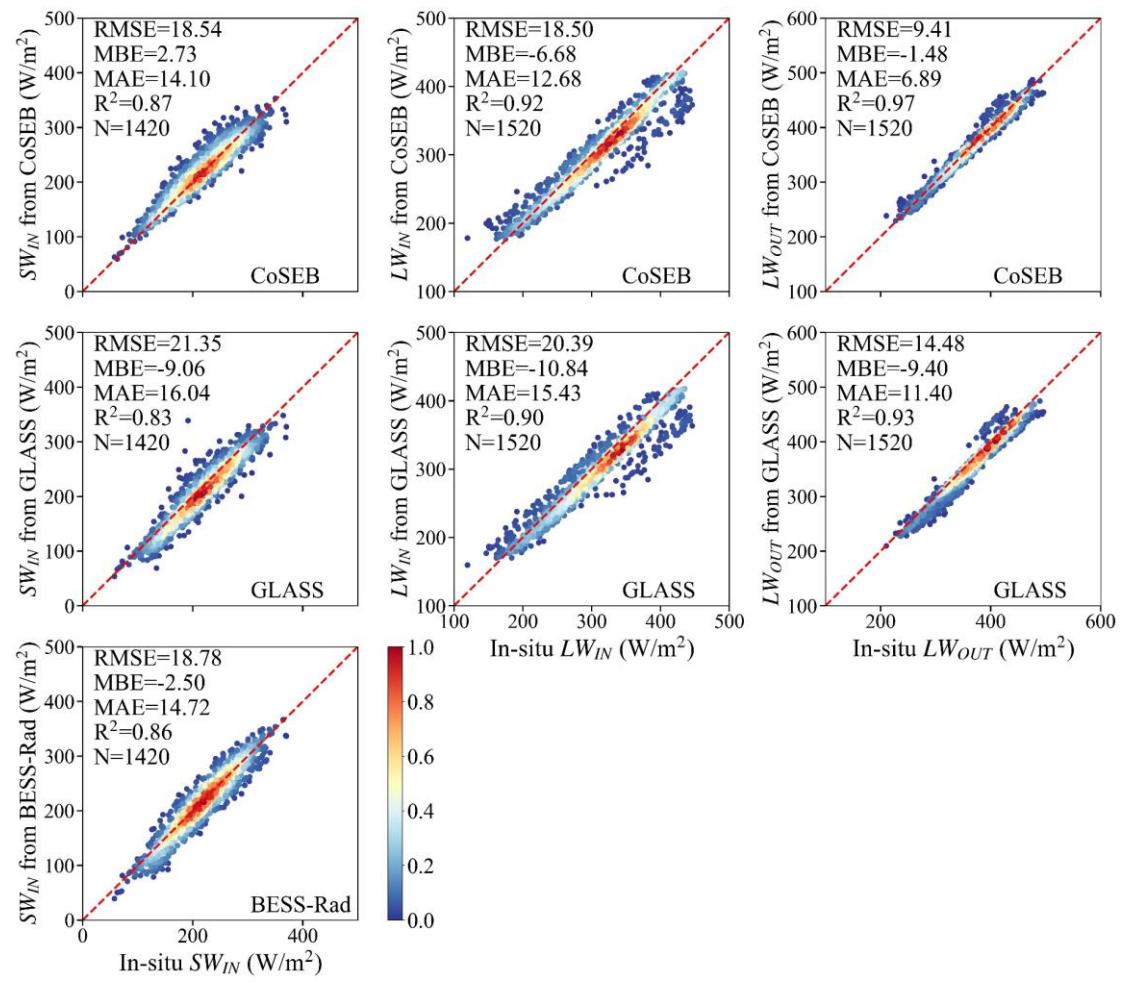
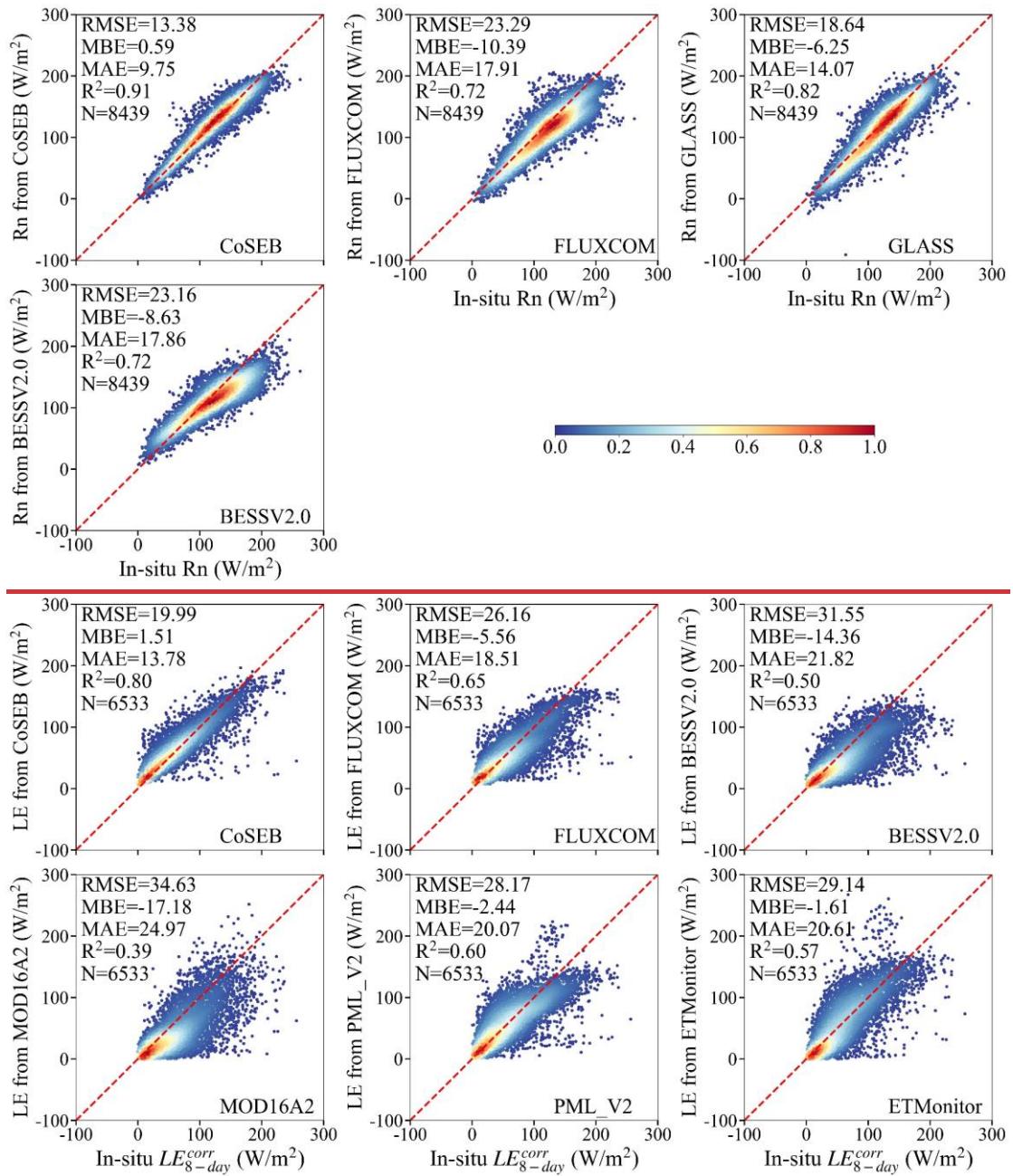
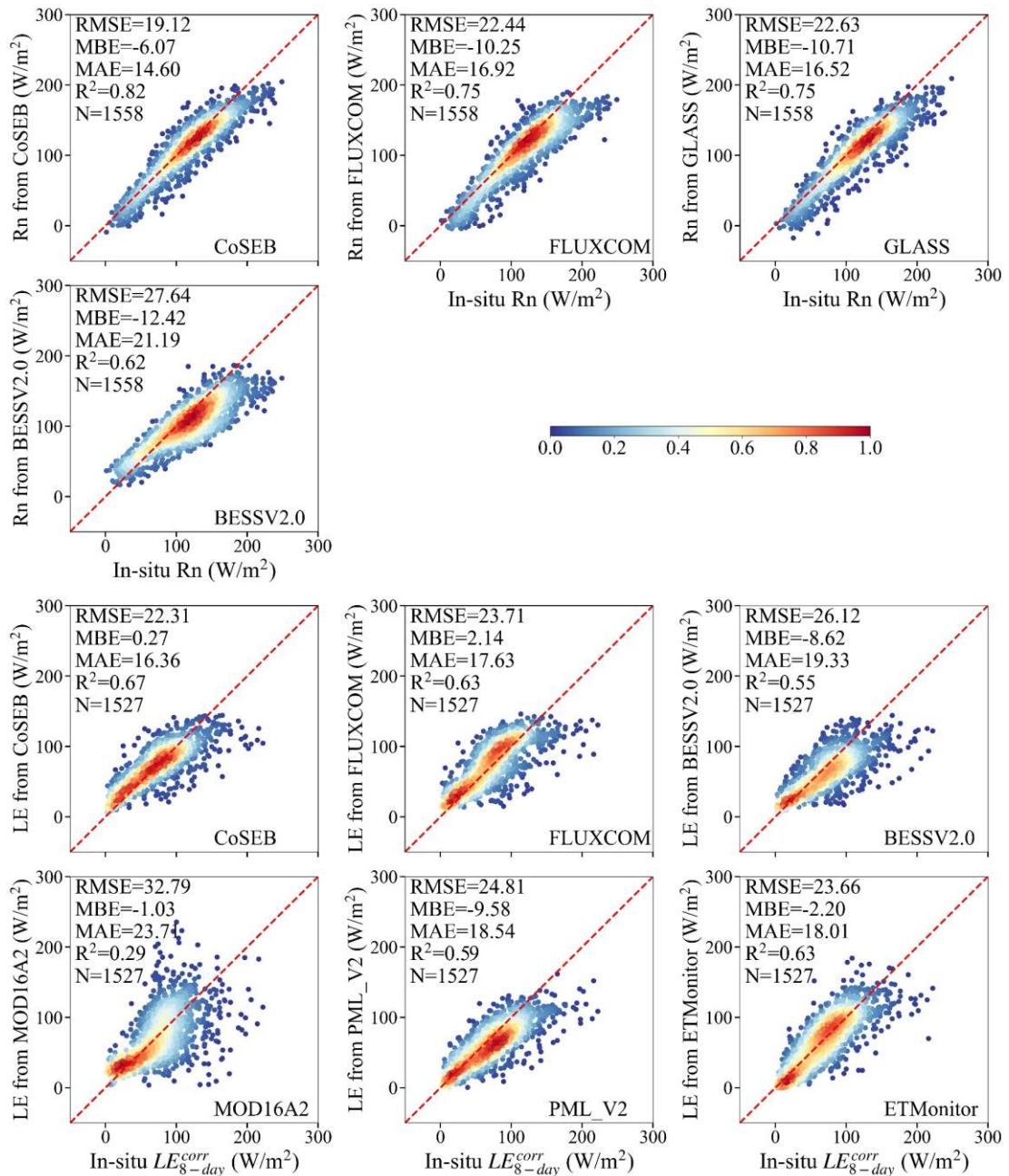
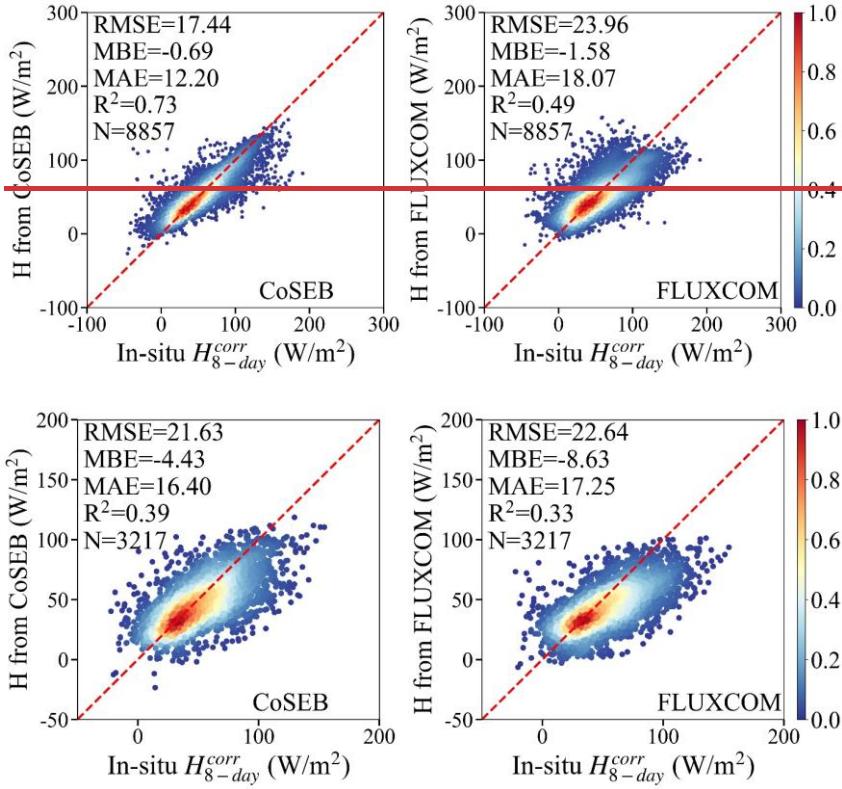


Fig. 8 Same as Fig. 6, but for the comparison at 8-day scale.





505  
506 **Fig. 9 Comparison of the 8-day net radiation (Rn, the upper two rows) and latent heat flux**  
507 **(LE, the lower three rows) from the CoSEB-based datasets, FLUXCOM, BESSV2.0, GLASS,**  
508 **MOD16A2, PML\_V2 and ETMonitor with in situ observed Rn, and energy imbalance-**  
509 **corrected LE (LE<sub>8-day</sub><sup>corr</sup>) at 134-44 tested sites. The colorbar represents the**  
510 **normalized density of data points.**



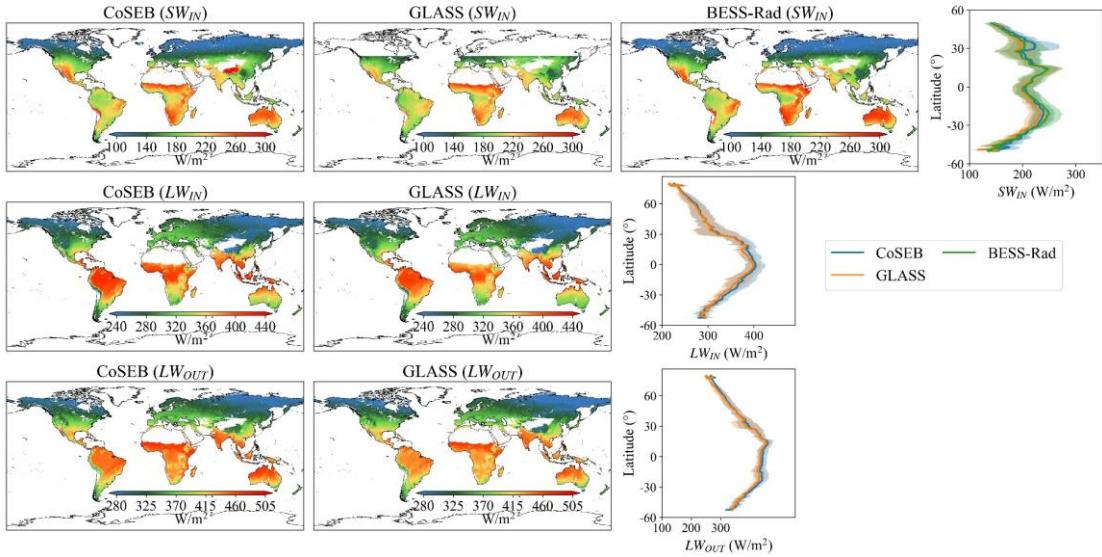
512  
513  
514 **Fig. 10 Comparison of the 8-day sensible heat flux (H) from the CoSEB-based datasets and**  
515 **the FLUXCOM with the in situ energy imbalance-corrected H ( $H_{8-day}^{corr}$ ) at 134-44 eddy**  
516 **covariance test sites. The colorbar represents the normalized density of data points.**

517 **4.3 Spatial-temporal patterns of global land surface radiation and heat fluxes**

518 In addition to the validation and inter-comparison of the CoSEB-based datasets at  
519 the global site scales, we further inter-compared the estimates of land surface radiation  
520 and heat fluxes from the CoSEB-based datasets and the mainstream products/datasets,  
521 in terms of their global spatial and temporal patterns.

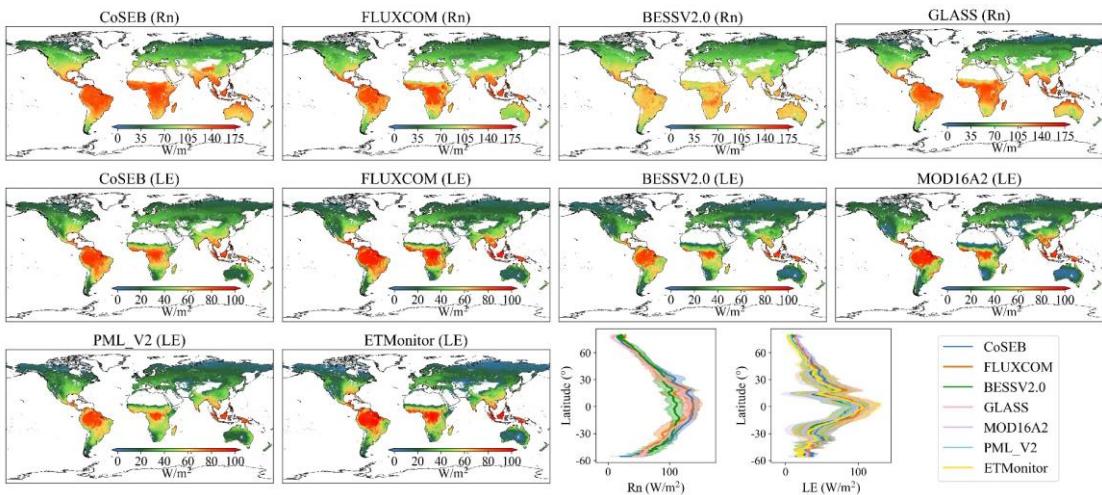
522 Figs. 11, 12 and 13 show the spatial distributions (excluding Greenland, Antarctic  
523 continent, deserts, water bodies and permanent snow) and latitudinal profiles of the  
524 global  $0.05^\circ$  mean annual  $SW_{IN}$ ,  $LW_{IN}$  and  $LW_{OUT}$ , Rn and LE, as well as H from 2001  
525 to 2018, respectively, as derived from the CoSEB-based datasets and mainstream  
526 products/datasets [i.e. GLASS, BESS-Rad, BESSV2.0, FLUXCOM, MOD16A2,  
527 PML\_V2 and ETMonitor, resampled to  $0.05^\circ$  using arithmetic averaging method or  
528 cubic convolutional method if necessary]. Overall, the spatial patterns of the estimates

529 from the CoSEB-based datasets aligned well with those observed in these mainstream  
530 products/datasets, though regional discrepancies were present. Specifically, the mean  
531 annual  $LW_{IN}$ ,  $LW_{OUT}$ , Rn, and LE generally exhibited decreasing trends from the equator  
532 towards higher latitudes, peaking in regions such as the Amazon Rainforest, Congo  
533 Rainforest, and the Malay Archipelago. In contrast, the higher mean annual  $SW_{IN}$  and  
534 H were mainly found in the Tibetan Plateau, southwestern U.S., mid-west Australia,  
535 Sahel and Southern Africa, while the lower values were found in high-latitude regions  
536 of  $>50^{\circ}\text{N}$ . In the region with of high values, the mean annual estimates of  $SW_{IN}$  from  
537 the CoSEB-based datasets were higher than those from GLASS but lower than those  
538 from BESS-Rad, the estimates of  $LW_{IN}$  and  $LW_{OUT}$  from the CoSEB-based datasets were  
539 both higher than those from GLASS, the estimates of Rn from the CoSEB-based  
540 datasets were significantly higher than those from BESSV2.0, and comparable to or  
541 slightly higher than those from FLUXCOM and GLASS, the estimates of LE from the  
542 CoSEB-based datasets were close to those from BESSV2.0 and PML\_V2, but slightly  
543 lower than those from FLUXCOM, MOD16A2 and ETMonitor. Besides, the estimates  
544 of H from the CoSEB-based datasets were higher than those from FLUXCOM in  
545 regions with high values, while lower than those from FLUXCOM in regions with low  
546 values.



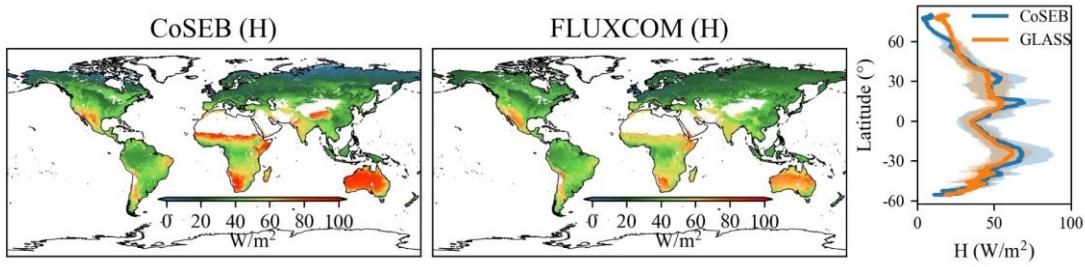
547

548 **Fig.11** Spatial patterns of global mean annual downward shortwave radiation ( $SW_{IN}$ , the first  
 549 row), downward longwave radiation ( $LW_{IN}$ , the second row) and upward longwave radiation  
 550 ( $LW_{OUT}$ , the third row) from 2001 to 2018 by CoSEB-based datasets, GLASS and BESS-Rad.  
 551 The rightmost subfigure of each row represents the latitudinal profiles of mean annual  $SW_{IN}$ ,  
 552  $LW_{IN}$  and  $LW_{OUT}$  from CoSEB-based datasets, GLASS and BESS-Rad, where the shaded area  
 553 represents the variation of standard deviation for each product.



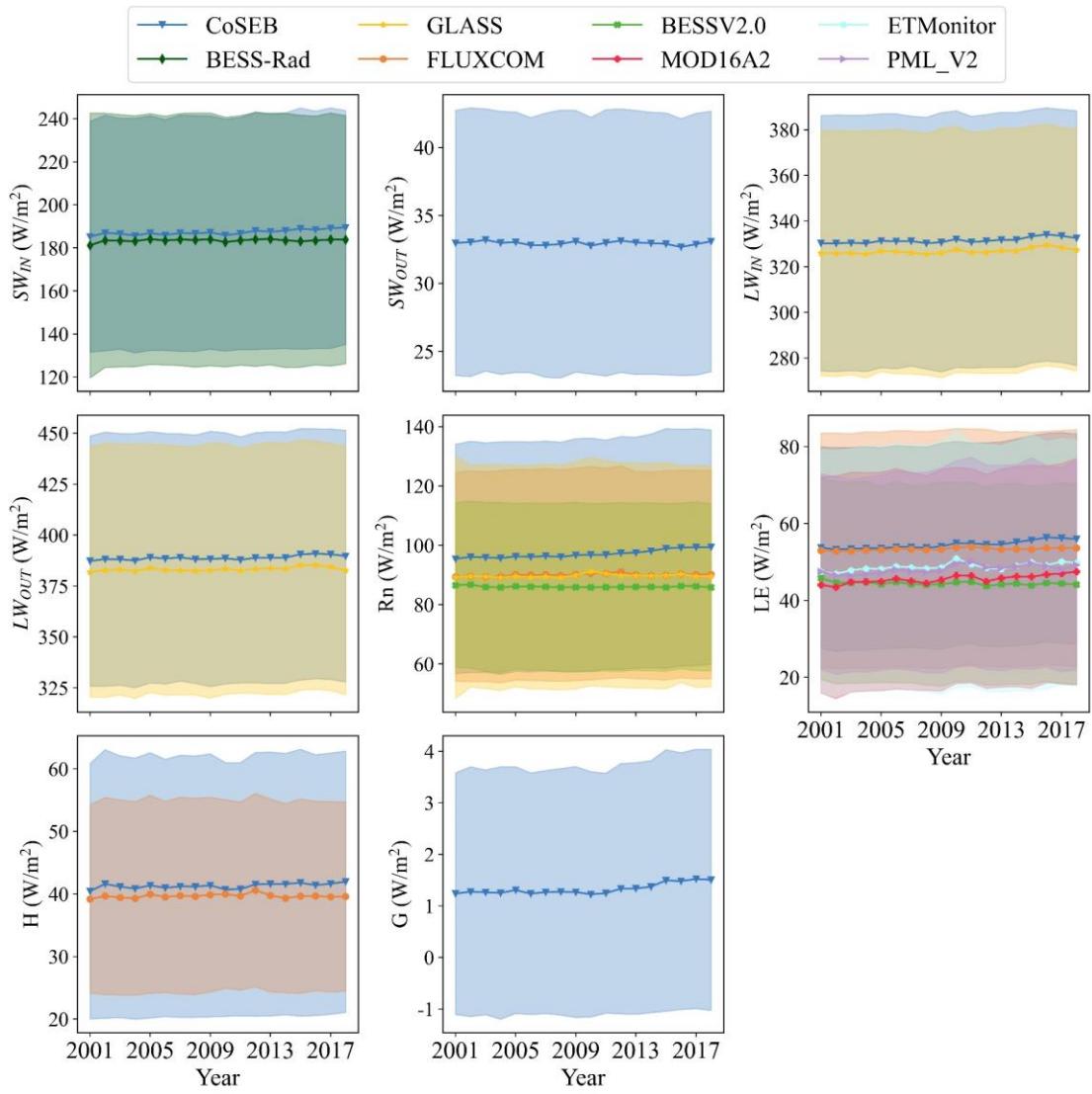
554

555 **Fig.12** Spatial patterns of global mean annual net radiation ( $Rn$ , the first row) and latent heat  
 556 flux ( $LE$ , the second and third rows) from 2001 to 2018 by CoSEB-based datasets, FLUXCOM,  
 557 BESSV2.0, MOD16A2, PML\_V2, ETMonitor and GLASS. The last two subfigures of the third  
 558 row represent the latitudinal profiles of mean annual  $Rn$  and  $LE$  from CoSEB-based datasets  
 559 and these mainstream products/datasets, where the shaded area represents the variation of  
 560 standard deviation for each product.



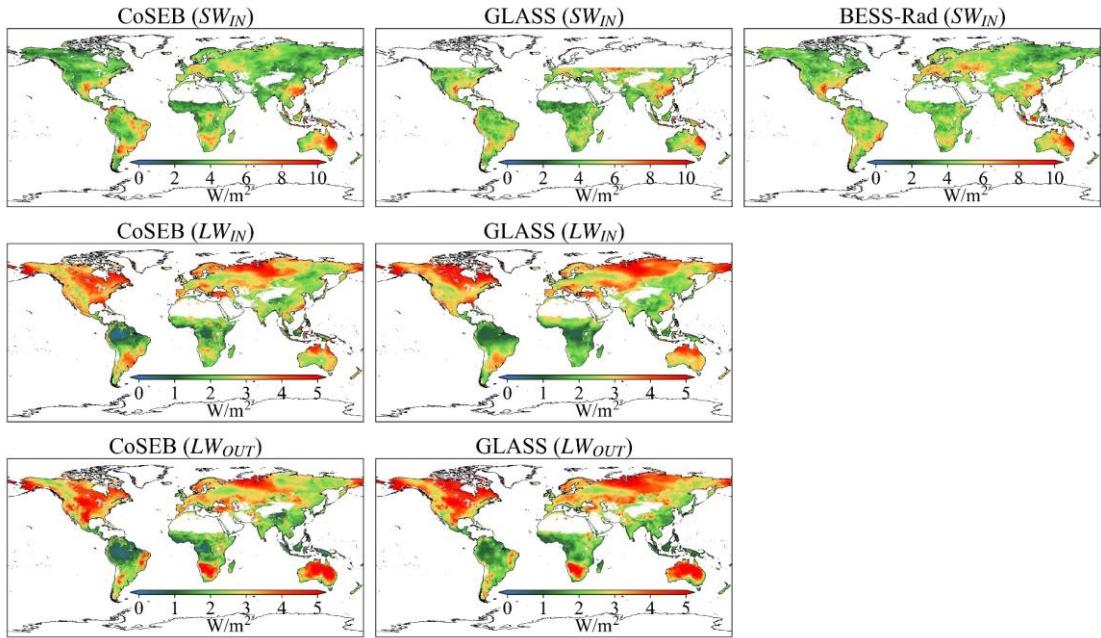
561  
 562 **Fig.13 Spatial patterns of global mean annual sensible heat flux ( $H$ ) from 2001 to 2018 by**  
 563 **CoSEB-based datasets and FLUXCOM. The rightmost subfigure represents the latitudinal**  
 564 **profiles of mean annual  $H$  from CoSEB-based datasets and FLUXCOM, where the shaded**  
 565 **area represents the variation of standard deviation for each product.**

566 The temporal evolutions of the global (excluding Greenland, Antarctic continent,  
 567 deserts, water bodies and permanent snow) land surface radiation and heat fluxes  
 568 derived from the CoSEB-based datasets and mainstream products/datasets from 2001  
 569 to 2018 were also investigated, as shown in Fig. 14. The results indicated that the  
 570 temporal variation of each flux from the CoSEB-based datasets generally agreed well  
 571 with those from mainstream products/datasets, exhibiting relatively stable trends. The  
 572 global annual mean estimates using area weighting average by the CoSEB-based  
 573 datasets from 2001 to 2018 varied between  $\sim 185.22$  and  $\sim 189.50$   $\text{W/m}^2$  with the mean  
 574 of  $\sim 187.23$   $\text{W/m}^2$  for  $SW_{IN}$ , between  $\sim 32.67$  and  $\sim 33.20$   $\text{W/m}^2$  with the mean of  $\sim 32.96$   
 575  $\text{W/m}^2$  for  $SW_{OUT}$ , between  $\sim 330.24$  and  $\sim 334.14$   $\text{W/m}^2$  with the mean of  $\sim 331.50$   $\text{W/m}^2$   
 576 for  $LW_{IN}$ , between  $\sim 387.25$  and  $\sim 390.82$   $\text{W/m}^2$  with the mean of  $\sim 388.81$   $\text{W/m}^2$  for  
 577  $LW_{OUT}$ , between  $\sim 95.41$  and  $\sim 99.39$   $\text{W/m}^2$  with the mean of  $97.11$   $\text{W/m}^2$  for  $Rn$ ,  
 578 between  $\sim 53.24$  and  $\sim 56.37$   $\text{W/m}^2$  with the mean of  $\sim 54.53$   $\text{W/m}^2$  for  $LE$ , between  
 579  $\sim 40.44$  and  $\sim 41.96$   $\text{W/m}^2$  with the mean of  $\sim 41.29$   $\text{W/m}^2$  for  $H$ , and between  $\sim 1.22$  and  
 580  $\sim 1.52$   $\text{W/m}^2$  with the mean of  $\sim 1.33$   $\text{W/m}^2$  for  $G$ . For each radiation or heat flux, the  
 581 annual mean estimates from the CoSEB-based datasets were overall higher than those  
 582 from the mainstream products/datasets. In particular, the annual mean  $Rn$  estimates  
 583 from the CoSEB-based datasets were higher than those from FLUXCOM, GLASS and  
 584 BESSV2.0 sequentially, and the annual mean  $LE$  estimates from the CoSEB-based  
 585 datasets were marginally higher than those from FLUXCOM, but substantially  
 586 exceeded those from ETMonitor, PML\_V2, MOD16A2 and BESSV2.0 sequentially.



587

588 **Fig. 14 Temporal variation of annual mean downward shortwave radiation ( $SW_{IN}$ ), upward**  
 589 **radiation ( $SW_{OUT}$ ), downward longwave radiation ( $LW_{IN}$ ), upward longwave**  
 590 **radiation ( $LW_{OUT}$ ), net radiation (Rn), latent heat flux (LE), sensible heat flux (H) and soil heat**  
 591 **flux (G) from 2001 to 2018 from the CoSEB-based datasets, BESS-Rad, GLASS, FLUXCOM,**  
 592 **BESSV2.0, PML\_V2, MOD16A2 and ETMonitor. The shaded area represents the variation of**  
 593 **standard deviation for each product.**

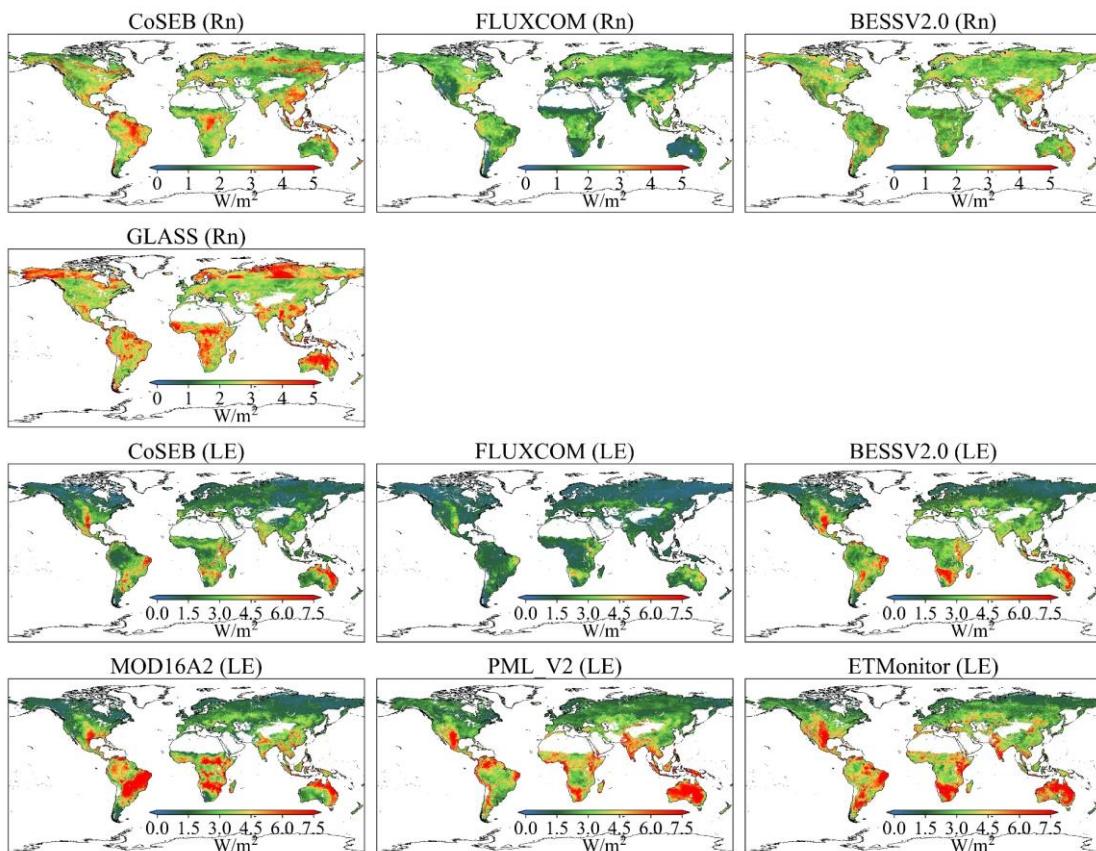


594

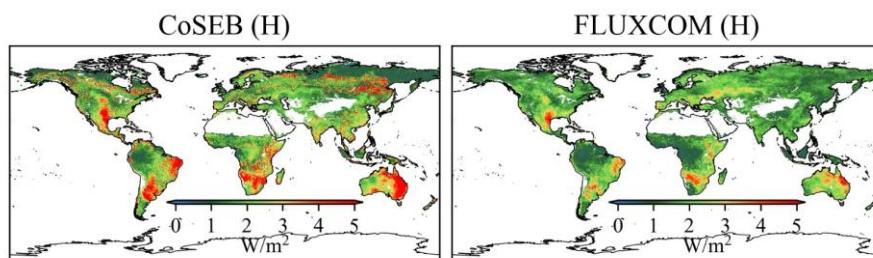
595 **Fig. 15 Spatial distribution of interannual variability (standard deviation) of downward**  
 596 **shortwave radiation ( $SW_{IN}$ , the first row), downward longwave radiation ( $LW_{IN}$ , the second**  
 597 **row) and upward longwave radiation ( $LW_{OUT}$ , the third row) from 2001 to 2018 by the CoSEB-**  
 598 **based datasets, GLASS and BESS-Rad.**

599 Figs. 15, 16 and 17 show the spatial patterns (excluding Greenland, Antarctic  
 600 continent, deserts, water bodies and permanent snow) of interannual variability of  $SW_{IN}$ ,  
 601  $LW_{IN}$  and  $LW_{OUT}$ , Rn and LE, as well as H from 2001 to 2018, respectively, derived  
 602 from the CoSEB-based datasets and mainstream products/datasets. In general, the  
 603 estimates from the CoSEB-based datasets displayed similar interannual variability in  
 604 space with those from the mainstream products/datasets. Specially, the estimates of  
 605  $SW_{IN}$  from the CoSEB-based datasets, BESS-Rad, and GLASS exhibited a significant  
 606 interannual variability mainly in northeastern Australia, eastern South America,  
 607 Southeast China, and Southwest North America. The interannual variability of  $LW_{IN}$   
 608 and  $LW_{OUT}$  by the CoSEB-based datasets and GLASS displayed high values primarily  
 609 at middle-to-high latitudes of the Northern Hemisphere and parts of Africa and  
 610 Australia. The interannual variability of Rn observed by the CoSEB-based datasets was  
 611 generally lower than that of GLASS, but higher than that of BESSV2.0 and FLUXCOM.  
 612 The CoSEB-based datasets missed the strong interannual variability of LE as observed  
 613 in MOD16A2, PML\_V2 and ETMonitor in parts of Africa, Australia and eastern South

614 America. Furthermore, FLUXCOM exhibited the weakest interannual variability of LE  
 615 in almost all regions. The interannual variability of H derived from the CoSEB-based  
 616 datasets was higher than ~~those that~~ from FLUXCOM, with stronger interannual  
 617 variabilities mainly observed in parts of eastern South America, southern Africa, and  
 618 northeastern Australia.



619  
 620 **Fig. 16 Spatial distribution of interannual variability (standard deviation) of net radiation (Rn,**  
 621 **the first and second rows) and latent heat flux (LE, the third and fourth row) from 2001 to**  
 622 **2018 by the CoSEB-based datasets, FLUXCOM, BESSV2.0, MOD16A2, PML\_V2,**  
 623 **ETMonitor and GLASS.**



624  
 625 **Fig. 17 Spatial distribution of interannual variability (standard deviation) of sensible heat flux**  
 626 **(H) from 2001 to 2018 by the CoSEB-based datasets and FLUXCOM.**

627 **5 Discussion**

628 Accurately monitoring the spatial and temporal variations of global land surface  
629 radiation and heat fluxes is crucial for quantifying the exchange of radiation, heat and  
630 water between the land and atmosphere under global climate change (Chen et al., 2020;  
631 Du et al., 2024; Kim et al., 2023; Liang et al., 2006; Wang et al., 2020). However,  
632 although numerous global RS-based products/datasets of land surface radiation and  
633 heat fluxes have been developed using physical and/or statistical methods, they  
634 typically provide either merely a single flux or multiple fluxes (see Table 1) that are  
635 estimated separately from uncoordinated models (Huang et al., 2024; Jung et al., 2019;  
636 Sun et al., 2023; Tang et al., 2019), leading to noticeable radiation imbalance and/or  
637 heat imbalance when these products are combined for practical applications. To address  
638 these limitations, we generated high-accuracy global datasets of land surface radiation  
639 and heat fluxes from 2000 to 2020 that adhere to both radiation and heat conservation  
640 laws, using our proposed CoSEB model (Wang et al., 2025).

641 Our CoSEB model, integrating underlying physical principles of training datasets  
642 into machine learning technique to effectively learn the interrelations among multiple  
643 targeted outputs, was originally designed for coordinating estimates of global land  
644 surface energy balance components (Rn, LE, H and G) to satisfy the energy  
645 conservation (Wang et al., 2025). Inspired by the idea of constructing the original  
646 CoSEB model, we further incorporated land surface radiation fluxes into our model to  
647 simultaneously consider the physical constraints of both surface radiation and heat  
648 conservation principles, by renewing the CoSEB using multiple remote sensing  
649 products and reanalysis datasets, as well as in-situ observations of  $SW_{IN}$ ,  $SW_{OUT}$ ,  $LW_{IN}$ ,  
650  $LW_{OUT}$ , Rn, LE, H and G. In selecting the 19 input variables to accommodate the  
651 additional target variables, prior knowledge derived from previous studies was  
652 employed to identify factors that exert significant influence on surface radiation and  
653 heat flux while maintaining relative inter-independence as much as possible (Jung et al.,  
654 2019; Mohan et al., 2020; Wang et al., 2021; Xian et al., 2024). This practice is

commonly adopted in data-driven models for estimating land surface water, energy, and carbon fluxes (Bai et al., 2024; Elghawi et al., 2023; Han et al., 2023; O. & Orth, 2021). The importance scores of the 19 different feature variables are exhibited in Table S4 in the Supplementary Material, and downward solar radiation, the primary source of the energy at the earth surface, is the most important input variable, consistent with the results from our previous study (Wang et al., 2025). Although some of the selected variables may exhibit a certain degree of multi-collinearity, each contributes unique and physically meaningful information, supporting the inclusion of all variables in model construction. To comprehensively account for the main factors influencing surface radiation and heat fluxes (Mohan et al., 2020; Wang et al., 2021; Xian et al., 2024) [JW1], the renewed CoSEB model utilized 19 easily accessible parameters/variables from ERA5 Land reanalysis datasets, GLASS products, MODIS products, GMTED2010 and NOAA/GML as input, which were readily available to generate datasets of global land surface radiation and heat fluxes in a practical and operational manner. (Wang et al., 2025) Note that the variable importance, derived from the built-in method of the random forests and potentially affected by multicollinearity among the input variables, is presented only as a reference. Retaining all 19 feature variables ensures the model's flexibility and generalization capability, enabling future incorporation of additional representative ground-based observations for further training and improvement. Besides, to investigate the impact of lagged effects of input variables on model performance, experiments were also conducted by adding lagged variables (e.g., the air temperature of the previous day) to the 19 input features. The results (Fig. S4 in the Supplementary Material) showed almost no improvement in model accuracy, suggesting that lagged effects on model performance were negligible within the CoSEB framework for estimates of daily surface radiation and heat fluxes. Furthermore, to better illustrate the effect of including additional radiation components ( $SW_{IN}$ ,  $SW_{OUT}$ ,  $LW_{IN}$  and  $LW_{OUT}$ ) in the renewed CoSEB model compared with the original version by Wang et al. (2025), we have tested the performance of a reconstructed model that

683 estimated only Rn, LE, H and G using the same independent variables and samples as  
684 those in the renewed CoSEB model. The results (Fig. S5 in the supplementary material)  
685 showed no significant differences in accuracy compared with those of the renewed  
686 CoSEB model, indicating the expansion of radiation components did not compromise  
687 model performance.

688 The main advantages of our CoSEB-based datasets of land surface radiation and  
689 heat fluxes lie in that [1] they are the first ~~RS-based~~data-driven global datasets that  
690 satisfy both surface radiation balance ( $SW_{IN} - SW_{OUT} + LW_{IN} - LW_{OUT} = Rn$   
691  $SW_{IN} - SW_{OUT} + LW_{IN} - LW_{OUT} = Rn$ ) and heat balance ( $LE + H + G = Rn$   
692  $LE + H + G = Rn$ ) among the eight fluxes, as demonstrated by both the RIR and EIR  
693 of 0, [2] the radiation and heat fluxes are characterized by high accuracies when  
694 validated against in-situ measurements at ~~134 “homogeneous”~~44 independent test sites  
695 (see the ~~first~~second paragraph in Section 4.22.1), where (1) the RMSEs for daily  
696 estimates of  $SW_{IN}$ ,  $SW_{OUT}$ ,  $LW_{IN}$ ,  $LW_{OUT}$ ,  $Rn$ ,  $LE$ ,  $H$  and  $G$  from the CoSEB-based  
697 datasets were 28.5137.52 W/m<sup>2</sup>, 10.394.20 W/m<sup>2</sup>, 14.2922.47 W/m<sup>2</sup>, 10.623.78 W/m<sup>2</sup>,  
698 22.409.66 W/m<sup>2</sup>, 24.3830.87 W/m<sup>2</sup>, 22.679.75 W/m<sup>2</sup> and 6.775.69 W/m<sup>2</sup>, respectively,  
699 as well as for 8-day estimates were 12.818.54 W/m<sup>2</sup>, 7.0812.19 W/m<sup>2</sup>, 9.2218.50 W/m<sup>2</sup>,  
700 8.349.41 W/m<sup>2</sup>, 13.389.12 W/m<sup>2</sup>, 19.9922.31 W/m<sup>2</sup>, 17.4421.63 W/m<sup>2</sup> and 4.254.60  
701 W/m<sup>2</sup>, respectively, (2) the CoSEB-based datasets, in comparison to the mainstream  
702 RS-based products/datasets (i.e. GLASS, BESS-Rad, FLUXCOM, BESSV2.0,  
703 MOD16A2, PML\_V2 and ETMonitor), better agreed with the in situ observations at  
704 ~~134 EC~~the 44 test sites, showing the RMSE reductions ranging from 4.350.01 W/m<sup>2</sup> to  
705 11.464.58 W/m<sup>2</sup> for  $SW_{IN}$ ,  $LW_{IN}$ ,  $LW_{OUT}$ ,  $Rn$  and  $LE$  at daily scale, and 4.620.24 W/m<sup>2</sup>  
706 to 14.640.48 W/m<sup>2</sup> for  $SW_{IN}$ ,  $LW_{IN}$ ,  $LW_{OUT}$ ,  $Rn$ ,  $LE$  and  $H$  at 8-day scale. Furthermore,  
707 the CoSEB-based datasets outperformed the ERA5-Land reanalysis datasets in  
708 estimating surface energy fluxes (where  $SW_{OUT}$ ,  $LW_{OUT}$ ,  $Rn$  and  $G$  for the ERA-Land  
709 were inferred from surface radiation balance and heat balance), particularly for  $SW_{OUT}$ ,  
710 H and G, with RMSE reductions of 0.13-8.15 W/m<sup>2</sup> when validated against in situ

711 observations at the 44 test sites (Figs. S6 and S7 in the Supplementary Material).  
712 Preliminary analysis indicates that the CoSEB-based datasets exhibit spatial patterns  
713 consistent with those of mainstream RS-based datasets and Earth system model outputs  
714 (see Fig. S8 in the supplementary material). More detailed analysis about their  
715 similarities and differences can be further conducted in future work.

716 Our developed datasets could be potentially applied in many fields, including but  
717 not limited to (1) exploring the spatial-temporal patterns of global land surface radiation  
718 and heat flux (es) and their driving mechanisms over the past decades under global  
719 change (e.g., rising CO<sub>2</sub> concentration, greening land surface and increasing air  
720 temperature), (2) investigating the variability of land surface radiation and heat fluxes  
721 caused by extreme events and human activities, e.g. afforestation or deforestation,  
722 wildfire, air pollution, weather extremes and urbanization, (3) assessing the resources  
723 of solar energy, geothermal energy, surface and ground water at regional and global  
724 scales, (4) monitoring natural hazards, e.g. drought in agriculture and forestry.

725 The uncertainties of our datasets are relevant to (1) the data preprocessing, and (2)  
726 the application of the CoSEB at model across different spatial scales. Specifically, the  
727 daily averages of surface radiation and heat fluxes for each day were obtained for  
728 analysis from good-quality half-hourly observations when the fraction of these good-  
729 quality half-hourly observations was greater than 80% in a day, due to the lack of  
730 consensus on the method for aggregating gapped half-hourly observations to daily data  
731 (Tang et al., 2024a; Yao et al., 2017; Zheng et al., 2022). Simple temporal interpolation  
732 of half-hourly in situ observations, which could therefore introduce substantial  
733 uncertainties, was not applied, because surface radiation and heat fluxes are sensitive  
734 to short-term variations in meteorological conditions and their intraday dynamics are  
735 often complex. Likewise, since there was no agreement on how to correct for the energy  
736 imbalance of turbulent heat fluxes, we adopted the most widely applied Bowen ratio  
737 method to enforce energy closure between  $Rn - G$  and  $LE + H$  (Castelli et al., 2018;  
738 Twine et al., 2000; Zhang et al., 2021). Another potential source of uncertainty arises

739 from differences in meteorological reanalysis data caused by spatial downscaling,  
740 which, as demonstrated in our previous study (Wang et al., 2025, the last paragraph of  
741 Section 5.1), has a relatively small impact on model estimates by the machine-learning-  
742 based CoSEB model combined with finer-resolution surface-related variables that  
743 partially compensate for the spatial heterogeneity and localized variations not captured  
744 by the coarse-resolution datasets. (Wang et al., 2025, the last paragraph of Section 5.1)  
745 (Wang et al., 2025, the last paragraph of Section 5.1) These data preprocessing had an  
746 effect on the construction of the renewed CoSEB model, which may further affect the  
747 global datasets. Moreover, the renewed CoSEB model was constructed at the spatial  
748 scale of 500 m to match the footprints of the in situ EC observations, but applied at the  
749 spatial resolution of 0.05° to generate global datasets, mainly limited by the computing  
750 and storage capabilities in-of our personal computers. However, the CoSEB-based  
751 datasets have also been validated and inter-compared at 134 EC44 independent test sites  
752 to demonstrate that the difference in spatial scale would not much affect the  
753 performance of the datasets. Despite these uncertainties, it is worth emphasizing that  
754 our work was the first attempt to innovatively develop data-driven energy-conservation  
755 datasets of global land surface radiation and heat fluxes with high accuracies.  
756

## 757 **6 Data availability**

758 The energy-conservation datasets of global land surface radiation and heat fluxes  
759 generated by the CoSEB model with spatial-temporal resolutions of daily and 0.05°  
760 from Feb.26, 2000 to Dec.31, 2020 are freely available through the National Tibetan  
761 Plateau Data Center at <https://doi.org/10.11888/Terre.tpdc.302559> (Tang et al., 2025a)  
762 and through the Science Data Bank (ScienceDB) at  
763 <https://doi.org/10.57760/sciencedb.27228> (Tang et al., 2025b).

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764 **7 Summary and Conclusion**

765 This study for the first time developed data-driven energy-conservation datasets  
766 of global land surface radiation and heat fluxes using our CoSEB model renewed based  
767 on GLASS and MODIS products, ERA5-Land reanalysis datasets, topographic data,  
768 CO<sub>2</sub> concentration data, and observations at 258 EC sites worldwide ~~from the~~  
769 ~~FLUXNET, AmeriFlux, EuroFlux, OzFlux, ChinaFLUX and TPDC~~.

770 The CoSEB-based datasets of land surface radiation and heat fluxes are the first  
771 ~~RS-based data-driven~~ global datasets that satisfy both surface radiation balance ( $SW_{IN} -$   
772  $SW_{OUT} + LW_{IN} - LW_{OUT} = Rn$ ) and heat balance ( $LE$   
773  $+ H + G = Rn$ ) among the eight fluxes. Meanwhile, the CoSEB-based  
774 datasets outperformed the mainstream products/datasets in accuracy. Specifically, at  
775 ~~134 44 EC independent test sites, the RMSEs (R<sup>2</sup>) for daily estimates of SW<sub>IN</sub>, SW<sub>OUT</sub>,~~  
776 ~~LW<sub>IN</sub>, LW<sub>OUT</sub>, Rn, LE, H and G from the CoSEB-based datasets were 37.52 W/m<sup>2</sup> (0.81),~~  
777 ~~14.20 W/m<sup>2</sup> (0.42), 22.47 W/m<sup>2</sup> (0.90), 13.78 W/m<sup>2</sup> (0.95), 29.66 W/m<sup>2</sup> (0.77), 30.87~~  
778 ~~W/m<sup>2</sup> (0.60), 29.75 W/m<sup>2</sup> (0.44) and 5.69 W/m<sup>2</sup> (0.44), respectively~~~~the RMSEs for~~  
779 ~~daily estimates of SW<sub>IN</sub>, SW<sub>OUT</sub>, LW<sub>IN</sub>, LW<sub>OUT</sub>, Rn, LE, H and G from the CoSEB-based~~  
780 ~~datasets were 28.51 W/m<sup>2</sup>, 10.39 W/m<sup>2</sup>, 14.29 W/m<sup>2</sup>, 10.62 W/m<sup>2</sup>, 22.40 W/m<sup>2</sup>, 24.38~~  
781 ~~W/m<sup>2</sup>, 22.67 W/m<sup>2</sup> and 6.77 W/m<sup>2</sup>, respectively~~, as well as for 8-day estimates were  
782 ~~12.81 18.54 W/m<sup>2</sup> (0.87), 7.08 12.19 W/m<sup>2</sup> (0.39), 9.22 18.50 W/m<sup>2</sup> (0.92), 8.34 9.41~~  
783 ~~W/m<sup>2</sup> (0.97), 13.38 9.12 W/m<sup>2</sup> (0.82), 19.99 22.31 W/m<sup>2</sup> (0.67), 17.44 21.63 W/m<sup>2</sup> (0.39)~~  
784 ~~and 4.25 4.60 W/m<sup>2</sup> (0.47), respectively~~. Moreover, the estimates from the CoSEB-  
785 based datasets in comparison to those from the mainstream products/datasets reduced  
786 the RMSE by ~~4.35~~ 0.01 W/m<sup>2</sup> to ~~11.46~~ 4.58 W/m<sup>2</sup> and increased the R<sup>2</sup> by ~~0.04~~ 0.01 to 0.3  
787 ~~09~~ for SW<sub>IN</sub>, LW<sub>IN</sub>, LW<sub>OUT</sub>, Rn and LE at daily scale, and reduced the RMSE by ~~4.62~~ 0.24  
788 W/m<sup>2</sup> to ~~14.64~~ 0.48 W/m<sup>2</sup> and increased the R<sup>2</sup> by ~~0.04~~ 0.01 to 0.41 38 for SW<sub>IN</sub>, LW<sub>IN</sub>,  
789 LW<sub>OUT</sub>, Rn, LE and H at 8-day scale, when these estimates were validated against in  
790 situ observations at ~~134 44 EC independent test~~ sites. Furthermore, the CoSEB-based  
791 datasets effectively captured the spatial-temporal variability of global land surface

792 radiation and heat fluxes, aligning well with those from the mainstream products.

793 Our developed datasets hold significant potential for application across diverse  
794 fields such as agriculture, forestry, hydrology, meteorology, ecology, and environmental  
795 science. They can facilitate comprehensive studies on the variability, impacts, responses,  
796 adaptation strategies, and mitigation measures of global and regional land surface  
797 radiation and heat fluxes under the influences of climate change and human activities.  
798 These datasets will provide valuable insights and data support for scientific research,  
799 policy-making, and environmental management, advancing global solutions to address  
800 climate change.

## 801 **Author contribution**

802 JW: Writing – original draft, Visualization, Software, Formal analysis, Data  
803 curation. RT: Writing – original draft, Validation, Supervision, Methodology, Funding  
804 acquisition, Formal analysis, Conceptualization. ML: Writing – review & editing,  
805 Validation. ZL: Writing – review & editing.

## 806 **Competing interests**

807 The authors declare that they have no conflict of interest.

## 808 **Acknowledgment**

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