

ABSTRACT

26 High-resolution three-dimensional (3D) wind field data are critical for a wide range of
27 applications, including wind energy assessment, low-altitude aviation, air quality modeling, and
28 extreme weather forecasting. Although ERA5 reanalysis remains widely used, its relatively coarse
29 spatial resolution (~31 km) limits its ability to capture local-scale atmospheric processes. To
30 address this, this study develops an hourly 3D dynamic wind field dataset with 1 km horizontal
31 resolution covering the Yangtze River Delta (YRD) region during the summer months (June–
32 August) from 2021 to 2023, namely YRD1km, generated through advanced dynamical
33 downscaling of ERA5 using a customized Weather Research and Forecasting (WRF) model
34 configuration. The methodology integrates multi-source observational nudging with high-
35 resolution land use parameterization to enhance near-surface wind accuracy and terrain-induced
36 flow representation, particularly in urban clusters and mountainous areas. Validation against
37 ground-based observations confirms the superior performance of YRD1km over ERA5 for hourly
38 10-m wind components, with Mean Absolute Error (MAE) reduced by 21.61% for U and 26.04%
39 for V, Root Mean Square Error (RMSE) reduced by 18.30% for U and 22.63% for V, and Nash–
40 Sutcliffe Efficiency (NSE) improved by 33.27% and 40.13%, respectively. On a daily mean basis,
41 both MAE and RMSE are reduced to below 0.5 m/s, and NSE reaches approximately 0.88.
42 Spatially, YRD1km captures finer spatial wind speed gradients and localized terrain-induced
43 circulations that are not captured by ERA5. Temporally, consistent accuracy improvements with
44 approximately 20% lower hourly error variability are seen when compared to ERA5. Vertically,
45 42.18% accuracy gains are observed in the near-surface layer when compared with radiosonde
46 profiles. Moreover, convective case analyses indicate that YRD1km captures vertically coherent
47 wind structures across multiple tropospheric levels that are closely linked to the initiation and
48 maintenance of deep convection, highlighting its diagnostic advantage in high-impact weather

49 events. Overall, the YRD1km 3D wind field dataset and its integrated methodological framework
50 provide a robust foundation for regional meteorological applications, including high-resolution
51 AI-based forecasting, renewable energy planning, and weather risk management in rapidly
52 developing regions such as the YRD. The YRD1km 3D wind field dataset is available at
53 <https://doi.org/10.57760/sciencedb.23752> (Zhang et al., 2025).

54 **Key words:** 3D wind field dataset; dynamical downscaling; multi-source observational
55 nudging; high-resolution land use; Yangtze River Delta

56

57 **1. Introduction**

58 Accurate characterization of three-dimensional (3D) wind fields with high spatiotemporal
59 resolution is fundamental to modern meteorological services, wind energy development, and the
60 safe operation of low-altitude economy. Although widely used ERA5 atmospheric reanalysis
61 datasets are capable of providing wind field variables that exhibit temporal continuity and physical
62 consistency, their relatively coarse spatial resolution limits the capability to resolve regional-scale
63 wind field features (Hu et al., 2023; Jung and Schindler, 2022), particularly in areas with complex
64 terrain and intense urbanization (Molina et al., 2021).

65 The Yangtze River Delta (YRD), as one of the most intensely urbanized regions in China,
66 exhibits evident spatiotemporal heterogeneity in local wind fields due to the combined effects of
67 sea-land thermal contrasts, urban heat island effects, and boundary layer turbulence (Zhang et al.,
68 2010). This presents significant challenges for precise wind energy resource assessment, urban
69 ventilation capacity diagnosis, and early warning of wind storm events. To address these
70 challenges, spatial downscaling of coarse-resolution reanalysis datasets has become a promising

71 strategy for improving regional wind field reanalysis and supporting fine-scale applications (Boé
72 et al., 2007; Tang et al., 2016; Zhang et al., 2020).

73 Spatial downscaling techniques primarily include statistical downscaling and dynamical
74 downscaling approaches. Statistical downscaling establishes statistical relationships between
75 coarse-resolution meteorological variables and local observational data (Dayon et al., 2015;
76 Tareghian and Rasmussen, 2013), enabling the acquisition of high-resolution wind field
77 information at relatively low computational costs (Zamo et al., 2016). However, such methods
78 often overlook the physical constraints among meteorological variables. In recent years, deep
79 learning has been increasingly applied to enhance the accuracy of statistical downscaling of wind
80 fields (Dujardin and Lehning, 2022; Dupuy et al., 2023; Höhle et al., 2020; Lian et al., 2024; Liu
81 et al., 2024a; Zhang and Li, 2021). Nevertheless, incorporating physical consistency into deep
82 learning frameworks remains a significant challenge (Sun et al., 2024). In contrast, dynamical
83 downscaling employs the fundamental equations governing the atmospheric dynamics to explicitly
84 resolve physical processes, thereby reconstructing regional weather systems at high resolutions
85 (Tang et al., 2016). Its effectiveness has been demonstrated in various applications (Bao et al.,
86 2015; Liu et al., 2024b; Xu et al., 2021). Horvath et al. (2012) applied the Weather Research and
87 Forecasting (WRF) model with sub-kilometer grid spacing over mountainous regions of Nevada
88 and showed that dynamical downscaling significantly improved the representation of near-surface
89 wind speed and variability compared to coarser reanalysis products. Notably, when combined with
90 nudging techniques, the model's responsiveness to the actual atmospheric state is further enhanced
91 (Harkey and Holloway, 2013; Lo et al., 2008).

92 Nudging, also known as Newtonian relaxation, is a data assimilation method that introduces
93 forcing terms into numerical model equations to incrementally adjust model variables toward

94 observations or analysis fields (Hoke and Anthes, 1976). Compared with variational assimilation
95 methods, nudging does not require the construction of an adjoint model or the estimation of
96 background error covariance matrices. As a result, it offers a simpler implementation and lower
97 computational cost (Daescu and Langland, 2013; Lei and Hacker, 2015). Research has
98 demonstrated that this method has been successfully applied in the construction of several high-
99 resolution reanalysis datasets. For example, the MERIDA HRES (4 km resolution, hourly) (Viterbo
100 et al., 2024) and the BAYWRF (1.5 km resolution, daily) (Collier and Mölg, 2020) datasets both
101 employ the WRF model to perform dynamical downscaling on ERA5 reanalysis data. By
102 integrating nudging techniques, these datasets have reconstructed local wind field characteristics
103 for Italy and the Bavarian region of Germany, respectively. Although dynamical downscaling
104 demands substantial computational resources, advancements in regional model structures and
105 high-performance computing technologies are expected to greatly improve its feasibility for
106 regional complex terrain studies and non-climate research applications (Gutowski et al., 2020;
107 Yuan et al., 2024).

108 Furthermore, accurate representation of land surface parameters is another critical factor
109 influencing the performance of wind field dynamical downscaling. In recent years, high-resolution
110 land use data have been increasingly incorporated into wind field modeling to optimize surface
111 parameterization (De Bode et al., 2023; Fu et al., 2020; Santos-Alamillos et al., 2015). The updated
112 land use datasets enable more precise characterization of various land surface features such as
113 urban areas, mountainous regions, and water bodies, which improve simulation of terrain-induced
114 flows and boundary layer processes, particularly in complex terrain regions (Golzio et al., 2021;
115 Siewert and Kroszczynski, 2023).

116 In summary, this study presents the development of a 1-km hourly 3D dynamic wind field
117 dataset over the YRD region (YRD1km), covering the period of the summer months (June to
118 August) from 2021 to 2023. The YRD1km dataset is generated by applying a state-of-the-art
119 dynamical downscaling technique to the ERA5 reanalysis data, integrating multi-source
120 observational nudging, and updating land surface information with high-resolution ESA
121 WorldCover 2020 (EWC2020) land use data. The resulting dataset provides enhanced accuracy in
122 simulating near-surface winds and tropospheric dynamic structures, particularly in urban and
123 mountainous areas where wind variability is often high.

124 This study evaluates the performance of YRD1km relative to ERA5, with a focus on both
125 horizontal and vertical wind field accuracy. It also assesses the effectiveness of an integrated
126 methodology that combines dynamical downscaling, observational nudging, and updated land use
127 data in improving wind field simulations over regions with complex land surface characteristics
128 and atmospheric variability. The findings highlight the potential of YRD1km to support a wide
129 range of applications, such as localized weather forecasting, renewable energy planning, air quality
130 modeling, and urban environmental management in rapidly urbanizing areas.

131 **2. Data**

132 **2.1 ERA5 Reanalysis Data**

133 The ERA5 reanalysis dataset (Hersbach et al., 2020), developed by the European Centre for
134 Medium-Range Weather Forecasts (ECMWF), integrates global multi-source observations
135 through 4D-Var data assimilation(<https://doi.org/10.24381/cds.bd0915c6>). It provides three-
136 dimensional hourly atmospheric variables (e.g., temperature, humidity, wind fields, and pressure)
137 with a horizontal resolution of $0.25^{\circ} \times 0.25^{\circ}$ (about 31km), serving as a widely adopted benchmark
138 in meteorological research. In this study, ERA5 supplies initial and boundary conditions for the

139 WRF model dynamical downscaling. Additionally, ERA5 serves as a baseline dataset for
140 comparative validation of YRD1km performance enhancements.

141 **2.2 Surface and Upper Air Weather Observations**

142 This study assimilates two observational datasets: (1) the NCEP ADP Global Upper Air and
143 Surface Weather Observations (<https://doi.org/10.5065/Z83F-N512>), comprising global terrestrial
144 stations, ocean buoys, ships, radiosondes, aircraft reports, and ASCAT satellite-derived winds from
145 the Global Telecommunication System (GTS), and (2) hourly data from Automatic Weather
146 Stations (AWS) operated by the China Meteorological Administration (CMA) (<http://data.cma.cn/>).
147 The spatial distributions of the two observational datasets over the YRD are illustrated in Figure
148 1a. The NCEP ADP dataset provides three-dimensional conventional meteorological
149 measurements from multiple observational platforms. As a complement to the NCEP ADP dataset,
150 the CMA AWS network delivers high-density surface observations across China, with a total of
151 2,169 stations—approximately six times the number of surface stations available from the NCEP
152 ADP dataset within the Chinese domain. This higher station density significantly enhances the
153 spatial representativeness of near-surface meteorological conditions in the YRD region. Using
154 Observation Nudging assimilation techniques, these datasets collectively correct systemic biases
155 in ERA5's near-surface fields within the WRF framework, enhancing the model's capacity to
156 resolve localized circulation patterns. The AWS data further act as a cross-validation source to
157 quantify YRD1km's accuracy improvements.

158 **2.3 High-resolution Land Cover Geographical Data**

159 Conventional land use datasets in WRF (USGS 1992-1993 or MODIS 2001) (Anderson et
160 al., 1976) are limited in their ability to reflect the rapid urban expansion and evolving land surface
161 characteristics of the YRD region. To address this, we integrate the EWC2020 dataset—a global

162 land cover product with 10-meter spatial resolution that classifies 11 surface types (e.g., built-up
163 areas, croplands, water bodies) (<https://esa-worldcover.org/en>). By updating WRF's land surface
164 parameters with EWC2020, we refine the representation of aerodynamic roughness lengths and
165 urban heat island effects. For instance, reclassifying Shanghai's Pudong district from Moderate
166 Resolution Imaging Spectroradiometer (MODIS) "mixed urban" to EWC2020 "high-intensity
167 built-up" improves wind field simulations by better capturing drag effects from high-rise structures,
168 as validated against AWS observations.

169 **3. Methods**

170 **3.1 WRF Model Configuration for Dynamical Downscaling**

171 This study employs the WRF-ARW model (v4.4.2) (Skamarock et al., 2019) to establish a
172 dynamic downscaling framework, enhancing the spatial resolution of ERA5 reanalysis data from
173 ~31 km to 1 km. The model domain is configured with a triple-nested grid centered at (29.36°N,
174 115.65°E) with horizontal resolutions of 9 km (D01, with 342×305 grid points), 3 km (D02, with
175 529×640 grid points), and 1 km (D03, with 919×949 grid points). The innermost domain, D03,
176 covers the entire YRD region (Figure 1a) and is designed to capture local circulation features
177 associated with urban clusters, lakes, and hilly terrain at a kilometer-scale resolution. **Two-way**
178 **nesting was employed for all three domains to allow feedback between nested grids. All domains**
179 **were initialized simultaneously, with ERA5 initial conditions interpolated directly onto each**
180 **domain, including the 1-km grid.** In the vertical direction, 61 terrain-following eta levels are used,
181 with the model top set at 10 hPa, which facilitates a detailed resolution of boundary layer dynamics.
182 **Based on previous studies (Bernini et al., 2025; Sahu et al., 2025),** the following physical
183 parameterization schemes were selected: the Thompson microphysics scheme (Thompson et al.,
184 2008), which is well-suited for high-resolution cloud microphysics; the Dudhia shortwave

185 radiation (Dudhia, 1989) and RRTM longwave radiation schemes (Mlawer et al., 1997) for
186 radiative transfer; and for boundary layer and land surface processes, the YSU non-local closure
187 scheme (Hong et al., 2006) coupled with the Noah land surface model (Tewari et al., 2004), which
188 enhances the representation of near-surface turbulent exchanges. The Kain-Fritsch cumulus
189 parameterization scheme (Kain, 2004) is applied only in the outer grid (D01) to mitigate
190 uncertainties in the “gray zone” below the 3 km grid resolution.

191 To reduce the accumulation of model errors, a cold-start strategy is implemented, with
192 simulations initiated four times daily at 00, 06, 12, and 18 UTC, respectively. Each run generates
193 a continuous 6-hour forecast period, from which the first hour is discarded as model spin-up.
194 Ultimately, this approach produces a continuous hourly three-dimensional wind field dataset.

195 **3.2 Conventional Observational Data Assimilation via Nudging**

196 While WRF dynamical downscaling enhances dataset resolution and preserves dynamical
197 constraints and physical consistency, it struggles to capture fine-scale wind field features over
198 complex underlying surfaces (e.g., urban clusters, water bodies) without dense observational
199 constraints. To address this, this study employs the Four-Dimensional Data Assimilation (FDDA)
200 technique, integrating conventional observations and ERA5 reanalysis fields through a Nudging
201 approach, thereby balancing localized dynamical processes and large-scale circulation consistency.

202 The core formulation of this approach is:

$$203 \quad \frac{\partial x}{\partial t} = F(x) + G \cdot W(t) \cdot (x_{obs} - x) \quad (1)$$

204 where x represents the model variable, $F(x)$ denotes the model dynamical equations, G is the
205 relaxation coefficient, and $W(t)$ is the temporal weighting function.

206 This study adopts a hybrid Nudging scheme combining two strategies: (1) Observation
207 Nudging (ON): Direct assimilation of in situ observations from CMA AWS and NCEP ADP to

208 dynamically refine local wind fields. (2) Analysis Nudging (AN): Application of ERA5 reanalysis
 209 fields as constraints to impose large-scale adjustments across the entire model domain hourly
 210 (Stauffer and Seaman, 1990), preventing deviations from large-scale circulation patterns. Thus, the
 211 combined ON+AN assimilation scheme ensures both large-scale consistency and enhanced
 212 regional meteorological representation.

213 Taking the nudging experiment conducted over a continuous four-day period from 1 to 4
 214 June 2022 as an illustrative example, the study quantitatively evaluates wind field accuracy over
 215 the YRD against ground-based observations using three statistical metrics: Mean Absolute Error
 216 (MAE), Root Mean Square Error (RMSE), and the Nash-Sutcliffe Efficiency coefficient (NSE;
 217 Nash and Sutcliffe, 1970), defined as follows:

$$218 \quad MAE = \frac{1}{n} \sum_{i=1}^n |A_i - O_i| \quad (2)$$

$$219 \quad RMSE = \sqrt{\frac{1}{n} \sum_{i=1}^n (A_i - O_i)^2} \quad (3)$$

$$220 \quad NSE = 1 - \frac{\sum_{i=1}^n (A_i - O_i)^2}{\sum_{i=1}^n (O_i - \bar{O})^2} \quad (4)$$

221 where A_i represents simulated values from either ERA5 reanalysis or the dynamically downscaled
 222 results, O_i denotes corresponding in situ observations, n is the total number of spatiotemporally
 223 matched observation–simulation pairs. The NSE metric ranges from $-\infty$ to 1, with values
 224 approaching 1 indicating perfect agreement between simulations and observations. As shown in
 225 Table 1, compared to ERA5 data, the ON+AN assimilated dataset demonstrates significant
 226 improvements across all statistical metrics for both the 10-m zonal (U10m) and meridional (V10m)
 227 wind components. In particular, the MAE is reduced by 22.26% for U10m and 27.45% for V10m,
 228 the RMSE is reduced by 18.38% for U10m and 23.31% for V10m, and the NSE is enhanced by
 229 33.27% for U10m and 41.27% for V10m. These results confirm that the ON+AN hybrid

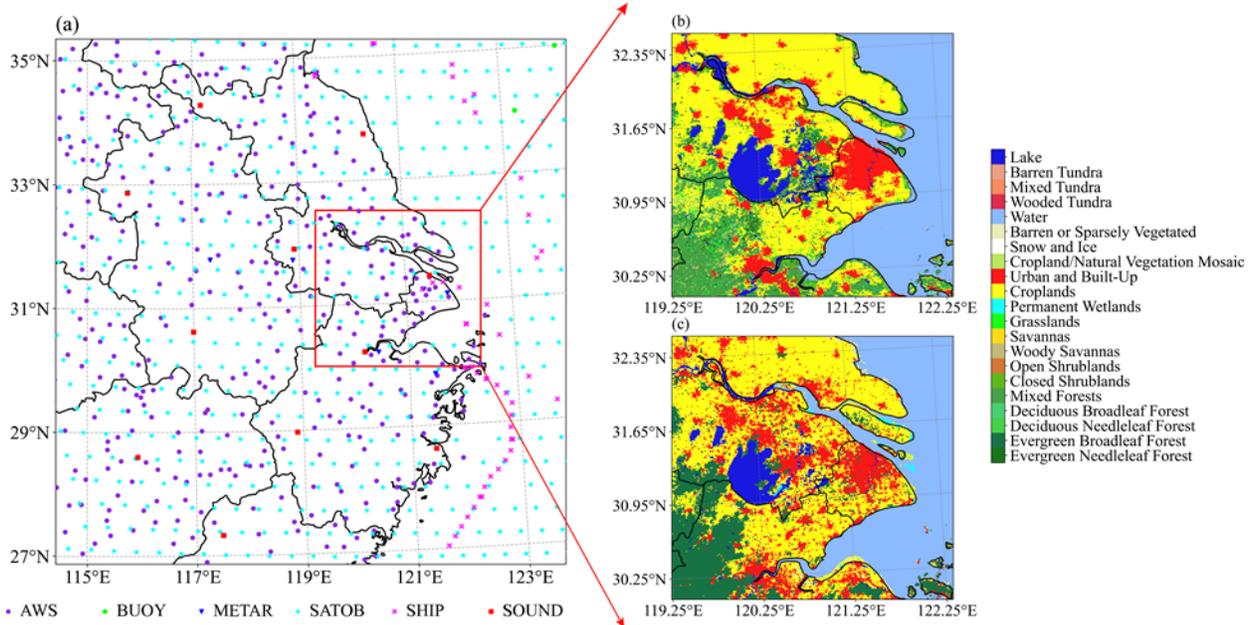
230 assimilation scheme substantially enhances the precision of high-resolution wind field datasets in
 231 the YRD region.

232 *Table 1. Comparison of surface (10-m) wind field performance between the ON+AN experiment*
 233 *and ERA5 reanalysis over the YRD region.*

Variable	Sample size	MAE (m/s)		RMSE (m/s)		NSE	
		ERA5	ON+AN	ERA5	ON+AN	ERA5	ON+AN
U10m	33172	1.213	0.943	1.583	1.292	0.504	0.669
V10m	33172	1.388	1.007	1.793	1.375	0.198	0.529

234 3.3 Impact of High-Resolution Land Use Data Updates

235 To address the impacts of rapid urbanization on wind field simulations in the YRD, this study
 236 enhances land surface characterization by updating the default MODIS 2001 land use data in the
 237 WRF model with the EWC2020 dataset at 10-meter resolution. Comparative analysis reveals
 238 substantial discrepancies between MODIS 2001 and EWC2020, particularly in Shanghai's
 239 metropolitan core (Figure 1b and 1c). The EWC2020 dataset resolves critical urban morphological
 240 features, including urban sprawl boundaries, park green spaces within city centers, and modified
 241 water-cropland interfaces, thereby more accurately capturing spatial heterogeneity in surface
 242 properties.



243

244 *Figure 1. Spatial distributions of key datasets used in this study. (a) Coverage of the innermost*
 245 *WRF domain (D03, 1-km resolution) over the YRD, along with the distribution of CMA Automatic*
 246 *Weather Stations (AWS) and the spatial coverage of NCEP ADP multi-source conventional*
 247 *observations used for nudging assimilation. The red box shows the region to highlight (b) Land*
 248 *use classification from the default MODIS 2001 dataset in WRF and (c) Updated high-resolution*
 249 *land use classification based on the EWC2020 product.*

250 To quantify land use update effects on wind field simulations, we conduct two experiments
 251 under the ON+AN assimilation framework: 1) LU-MODIS: Retains default MODIS-based land
 252 use types; 2) LU-EWC2020: Incorporates the refined EWC2020-derived surface parameters.
 253 Using the four-day case study, validation metrics (Table 2) demonstrate small but obvious positive
 254 impacts across all metrics for the LU-EWC2020 experiment compared to LU-MODIS. These
 255 results confirm the value of high-resolution land use updates in resolving urbanization-induced
 256 land-atmosphere interactions.

257 *Table 2. Statistical evaluation of land use sensitivity experiments conducted over the YRD region.*

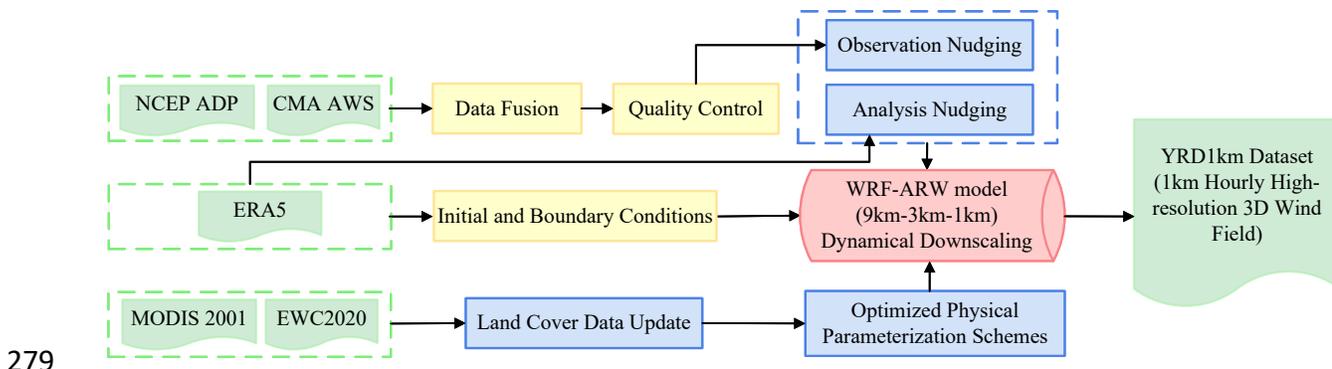
Variable	Sample size	MAE (m/s)		RMSE (m/s)		NSE	
		LU-MODIS	LU-EWC2020	LU-MODIS	LU-EWC2020	LU-MODIS	LU-EWC2020
U10m	33172	0.943	0.935	1.292	1.278	0.669	0.676
V10m	33172	1.007	0.992	1.375	1.355	0.529	0.542

258 **3.4 High-Resolution 3D Wind Field Dataset Generation**

259 Building on the evaluation results in section 3.2 and 3.3, this study develops a systematic
 260 framework for generating the YRD1km dataset over the YRD region, as shown in Figure 2. In the
 261 preprocessing stage, observational constraints for nudging were derived from the integration and
 262 quality control (QC) of NCEP ADP and CMA AWS datasets. Surface parameterization was refined
 263 by replacing the default MODIS 2001 land-use data with the updated **EWC 2020** dataset. For
 264 model simulation, ERA5 reanalysis provided the initial and boundary conditions for a triple-nested
 265 WRF configuration (9 km → 3 km → 1 km). The updated surface parameters were used to
 266 optimize the static fields, while a suite of optimized physical schemes and a cold-start initialization
 267 strategy were applied to suppress error accumulation. A hybrid observational nudging scheme (ON
 268 + AN) was employed to enhance the model’s consistency with observed atmospheric states,
 269 resulting in continuous hourly 3D wind vector outputs at 1-km horizontal resolution and 61 vertical
 270 levels during the summer months (June – August) from 2021 to 2023.

271 Comprehensive multi-dimensional validation was performed using both surface station
 272 observations and radiosonde profiles. The near-surface wind simulation performance was assessed
 273 through MAE, RMSE, and NSE metrics, to evaluate the overall, spatial, and temporal accuracy of
 274 the dataset. In addition, radiosonde-derived vertical wind profiles were used to examine the fidelity
 275 of the reconstructed wind field structure in the lower and middle troposphere. Furthermore,
 276 **convective case analyses** highlight the capability of the YRD1km dataset to capture fine-scale

277 dynamical features, demonstrating clear improvements over ERA5 and underscoring the
 278 effectiveness of the integrated approach in high-resolution wind field reconstruction.



279

280 *Figure 2. The schematic workflow of YRD1km 3D wind field generation.*

281 4. Results and Discussion

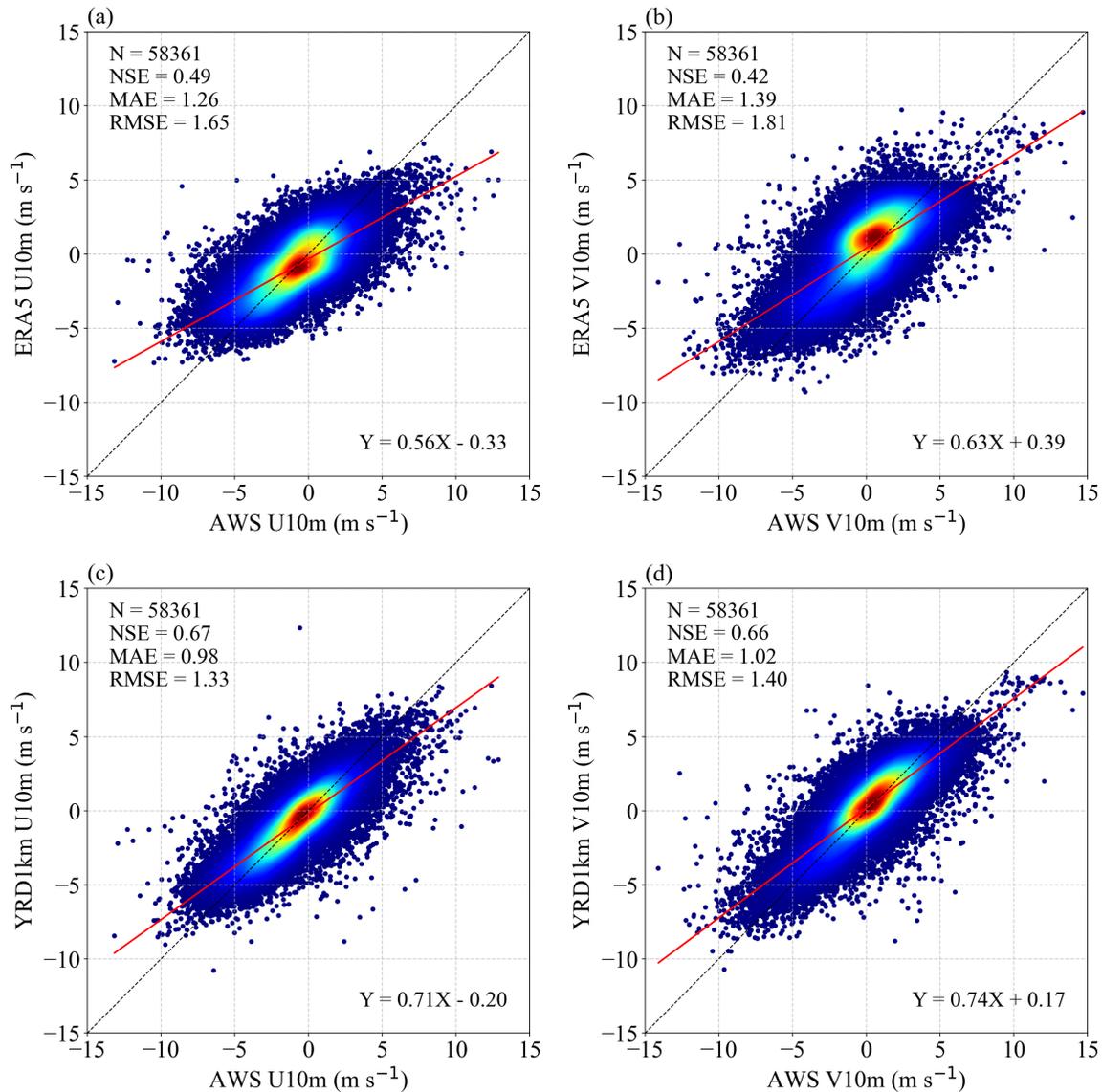
282 4.1 Evaluation of YRD1km High-Resolution Dataset Accuracy

283 4.1.1 Accuracy Evaluation of YRD1km and ERA5 Based on AWS Observations

284 **A comprehensive evaluation of near-surface wind field accuracy was conducted using AWS**
 285 **observations across the YRD region over a one-week period from 1 to 7 June 2022, serving as a**
 286 **representative example.** Due to the different spatial resolutions of YRD1km and ERA5, a nearest-
 287 grid-point matching method was adopted for comparison with station observations (Liu et al.,
 288 2025). As shown in Figure 3, scatterplots of the 10-m wind field U and V components for both
 289 ERA5 and YRD1km datasets were analyzed to assess their respective simulation capabilities.
 290 Overall, YRD1km exhibited superior performance in both U and V components, as evidenced by
 291 higher NSE coefficients, lower MAE and RMSE, and a tighter scatter distribution. Regression
 292 slopes for YRD1km were also notably closer to the 1:1 reference line, indicating a more accurate
 293 representation of the near-surface wind field compared to ERA5. For the U component (Figure 3a,
 294 c), ERA5 presented an NSE of **0.49**, with MAE and RMSE of **1.26 m/s** and **1.65 m/s**, respectively,
 295 and a regression slope of only **0.56**, with increasing deviations under higher wind speed conditions.

296 In contrast, YRD1km achieved a significant improvement with an NSE of **0.67**, MAE reduced to
297 **0.98** m/s, RMSE reduced to **1.33** m/s, and an increased regression slope of **0.71**, significantly
298 reducing systematic biases. Further analysis based on the sign of the U component revealed that
299 ERA5 exhibited a consistent underestimation of both easterly winds ($U < 0$) and westerly winds
300 ($U > 0$), particularly under stronger wind conditions ($|U| > 2$ m/s). This finding aligns with previous
301 reports by Hu et al. (2023). While YRD1km also exhibited a similar underestimation pattern, its
302 magnitude was notably reduced, indicating an improved representation of directional wind
303 components compared to ERA5. Additionally, as wind speed increased, scatter dispersion became
304 more pronounced, with fewer samples in the high wind speed range, adding challenges to accurate
305 simulation.

306 For the V component (Figure 3b, d), ERA5 showed an even lower NSE of **0.42**, with MAE
307 and RMSE of **1.39** m/s and **1.81** m/s, respectively, and a regression slope of **0.63**, indicating a less
308 accurate simulation. Conversely, YRD1km significantly improved the NSE to **0.66**, reduced MAE
309 to **1.02** m/s, RMSE to **1.40** m/s, and increased the regression slope to **0.74**. Similar to the U
310 component, the V component displayed a directional-dependent error pattern, with an
311 underestimation of both northerly winds ($V < 0$) and southerly winds ($V > 0$), especially under
312 stronger wind conditions. The increasing scatter dispersion and simulation uncertainty with higher
313 wind speeds further highlight the challenges and needs of reproducing complex wind fields.



314

315 *Figure 3. Scatterplot evaluation of 10-m wind components over the YRD region: (a) ERA5*

316

U10m, (b) ERA5 V10m, (c) YRD1km U10m, and (d) YRD1km V10m.

317

Results in Figure 3 are based on hourly data. Considering that climate research emphasizes

318

the use of daily data to smooth short-term fluctuations and reveal long-term trends (Kotlarski et

319

al., 2019; Nashwan et al., 2019; Zhang et al., 2024), this study further examined the simulation

320

accuracy of 10-m wind filed at the daily mean scale. **Based on daily mean observations from more**

321

than 300 AWS stations across the YRD region over a continuous seven-day period from 1 to 7

322 June 2022 (Table 3), the comparison results demonstrate that YRD1km maintains a consistent
 323 accuracy advantage over ERA5 for all evaluated metrics, including the U and V wind components
 324 as well as 10-m wind speed. Notably, the daily mean values of the U and V components exhibited
 325 better statistical performance than their hourly counterparts, as temporal averaging effectively
 326 mitigates short-term fluctuations and random errors, enhancing simulation stability. Additionally,
 327 compared to 10-m wind speed (WSPD10m), the U and V components demonstrated greater
 328 improvements in error metrics, with NSE values closer to 1. This is primarily because wind speed
 329 is a scalar variable, while U and V components are vectors accounting for wind direction errors.

330 *Table 3. Statistical comparison of daily 10-m wind fields between ERA5 and YRD1km datasets*
 331 *over the YRD region.*

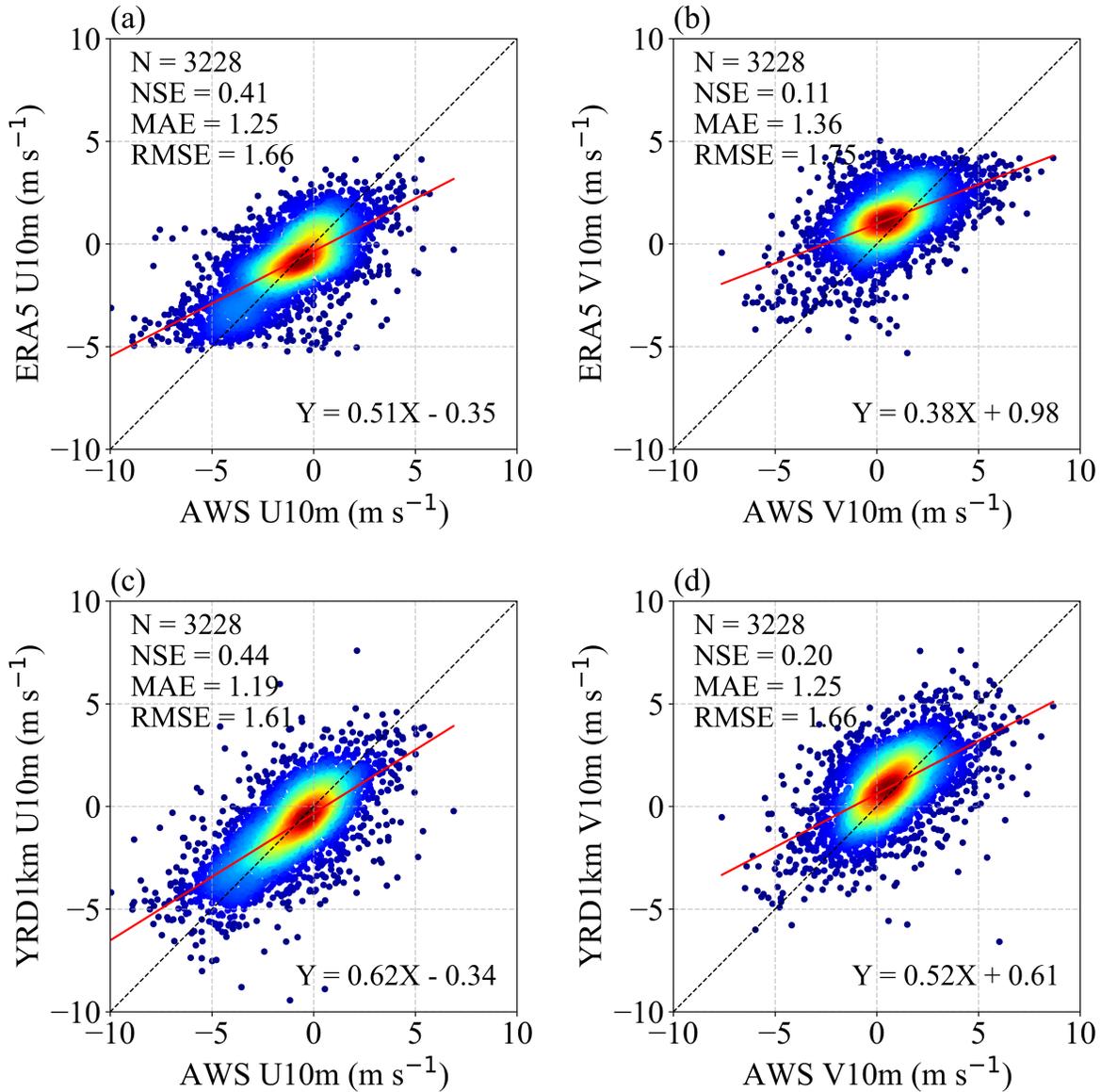
Variable	Indicator	Data		Improvement (%)
		ERA5	YRD1km	
U10m	MAE (m/s)	0.598	0.338	43.48
	RMSE (m/s)	0.783	0.451	42.40
	NSE	0.753	0.918	66.80
V10m	MAE (m/s)	0.716	0.367	48.74
	RMSE (m/s)	0.946	0.498	47.36
	NSE	0.583	0.884	72.18
WSPD10m	MAE (m/s)	0.619	0.422	31.83
	RMSE (m/s)	0.820	0.547	33.29
	NSE	0.470	0.765	55.66

332 **4.1.2 Independent Validation**

333 To rigorously evaluate the generalization capability of the YRD1km dataset beyond locations
 334 directly constrained by the nudging procedure, we implemented an independent validation strategy
 335 using a subset of AWS stations withheld from the assimilation process. Specifically, approximately

336 10% of the stations employed in the ON+AN nudging experiments were randomly selected and
337 reserved exclusively for independent evaluation, while the remaining stations continued to provide
338 observational constraints for WRF simulations.

339 The independent validation was conducted over a continuous four-day period from 1 to 4
340 June 2022, allowing the evaluation to sample a range of synoptic and mesoscale meteorological
341 conditions. Validation against observations at these withheld stations shows that the YRD1km
342 dataset exhibits improved performance relative to ERA5 in simulating near-surface wind
343 fields. For the U component, NSE increases by 5.08%, MAE decreases by 4.80%, and RMSE
344 decreases by 3.01% relative to ERA5. For the V component, NSE increases by 10.11%, while
345 MAE and RMSE decrease by 8.09% and 5.14%, respectively. In addition, the fitted relationships
346 between simulated and observed U and V components shifted closer to the 1:1 line, indicating
347 enhanced fidelity in reproducing near-surface wind variability (Figure 4). These results indicate
348 that the performance gains in YRD1km are not limited to assimilated locations. Instead, the
349 improvements extend spatially, demonstrating that the combination of dynamical downscaling,
350 high-resolution land-surface updates, and the hybrid ON+AN nudging framework enhances the
351 regional wind field representation in a physically coherent manner.



352

353 *Figure 4. Independent validation scatterplots of 10-m wind components over the YRD region: (a)*

354

ERA5 U10m, (b) ERA5 V10m, (c) YRD1km U10m, and (d) YRD1km V10m.

355 **4.1.3 Comparison of spatial variations between YRD1km and ERA5**

356

Building upon the preceding quantitative accuracy assessment, the study further examines

357

the spatial variations of near-surface wind fields represented by the YRD1km and ERA5 datasets,

358

as illustrated in Figure 5. Overall, while both datasets (Figure 5a and 5c) adequately capture the

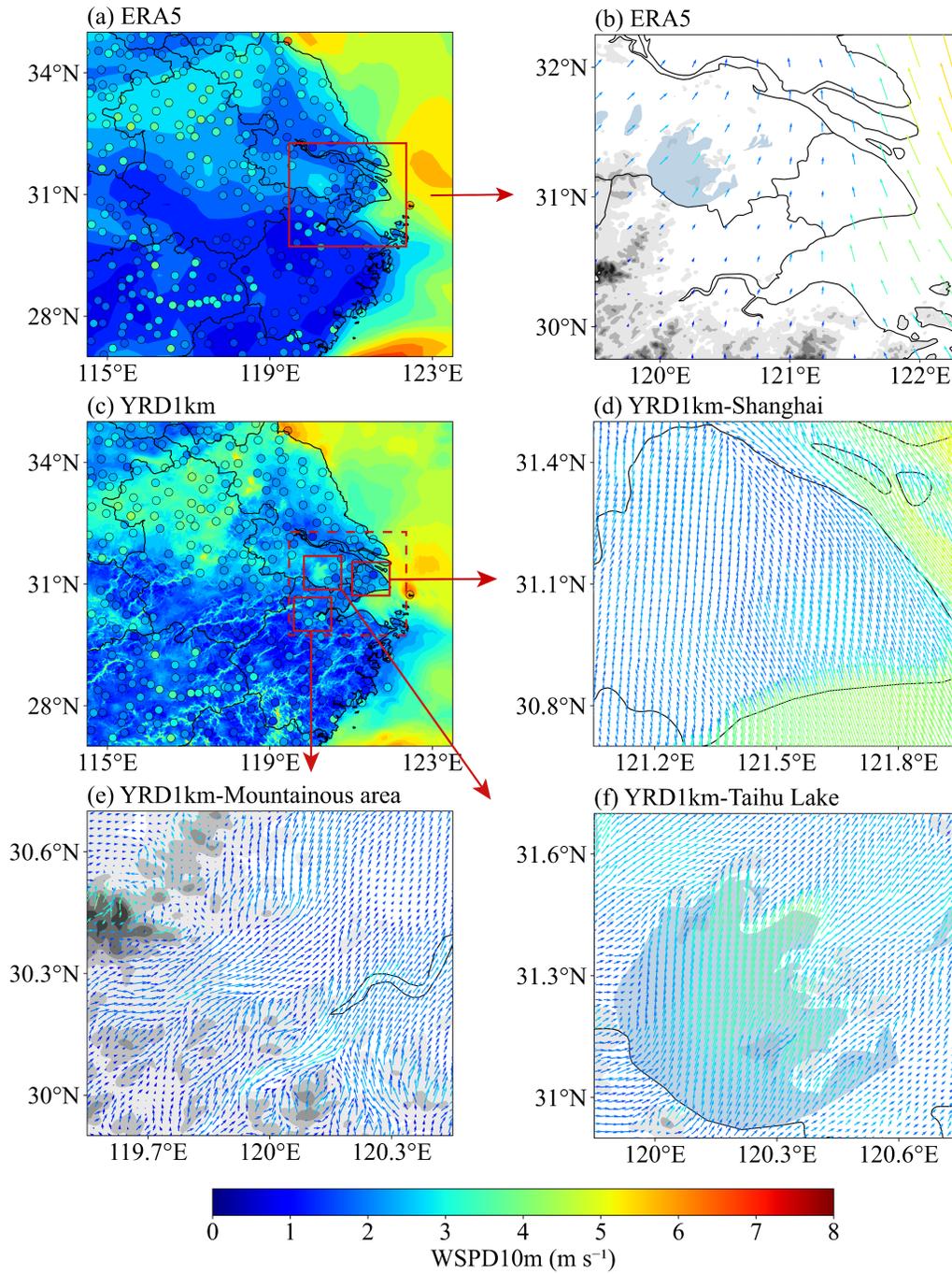
359

large-scale spatial variations of 10-m wind speeds across the YRD, YRD1km demonstrates a

360 notable advantage in resolving mesoscale and local-scale wind field characteristics. Specifically,
361 YRD1km (Figure 5c) offers a much finer spatial representation of wind speed variations compared
362 to ERA5, closely aligned with observational data, particularly over complex terrain and urbanized
363 areas. This includes enhanced wind speed zones over large water bodies such as Lake Taihu,
364 realistic gradients in mountainous regions like southern Anhui and Zhejiang driven by valley flows
365 and orographic effects, as well as improved wind speed structures over highly urbanized areas such
366 as Shanghai. Furthermore, ERA5 exhibits underestimation of wind speed maxima near offshore
367 observation points (e.g., in the East China Sea). YRD1km mitigates these biases through
368 assimilation of AWS data via a nudging approach, enabling better alignment with ground truth
369 observations and significantly enhancing the fidelity of simulated wind fields.

370 These spatial advantages are further highlighted through detailed analyses of wind vector
371 fields. As shown in Figure 5b, ERA5 exhibits an overly smoothed wind field with limited flow
372 differentiation near topographic boundaries. In contrast, the YRD1km dataset presents highly
373 structured and terrain-conforming wind directions. Over the Shanghai metropolitan area (Figure
374 5d), the wind field aligns with urban morphological structures, showing clear directional deflection
375 near city boundaries and dense river network regions, primarily due to thermal forcing and surface
376 drag associated with urbanization. In the mountainous region near Hangzhou (Figure 5e), the wind
377 field captures pronounced curvature and flow separation that closely follow terrain contours,
378 effectively representing multiple terrain-induced processes such as valley and slope winds. Over
379 Lake Taihu (Figure 5f), YRD1km simulates a divergent wind pattern, with significantly higher
380 wind speeds over the lake surface relative to surrounding land, indicative of thermally driven lake–
381 land breeze circulations.

382 Collectively, the spatial patterns observed in both scalar (wind speed) and vector (wind
383 direction) fields strongly affirm the capability of YRD1km to resolve sub-regional atmospheric
384 dynamics. These results further highlight the dataset's potential for supporting a broad spectrum
385 of regional meteorological applications.



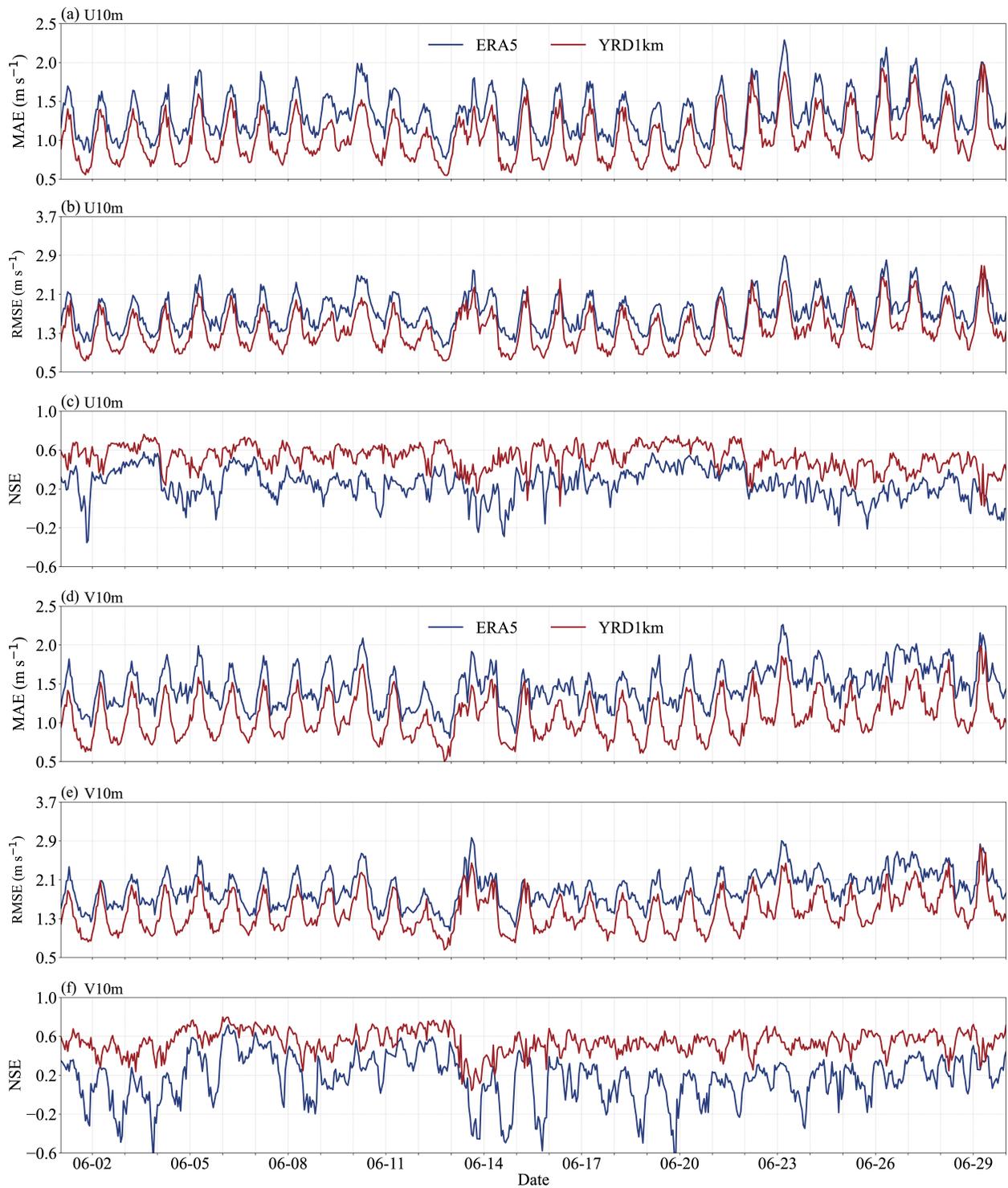
386

387 *Figure 5. Spatial distribution of daily mean near-surface wind fields over the YRD region on 1*
388 *June 2022. Panels (a) and (c) show daily mean 10-m wind speed (WSPD10m) from the ERA5 and*
389 *YRD1km datasets, respectively, overlaid with AWS station observations (colored dots). Panels (b),*
390 *(d), (e), and (f) show locally enlarged wind vector fields: (b) ERA5 over Shanghai and its*
391 *surrounding urban agglomeration; (d) YRD1km over the Shanghai metropolitan area; (e) the*
392 *mountainous region near Hangzhou; and (f) Lake Taihu. Arrows are color-coded by wind speed*
393 *magnitude and overlaid on shaded terrain elevation, with darker tones indicating higher altitudes.*
394 *Major water bodies are shaded in steelblue for clearer identification.*

395 **4.2 Statistical Analysis of the Long-term Time Series of Surface Wind**

396 To assess the temporal performance of the proposed YRD1km dataset, hourly time series
397 analyses of the U10m and V10m wind components were conducted over the YRD region for June
398 2022. For each hour, the performance metrics were computed by spatially averaging errors over
399 all available surface stations within the region. Figures 6 presents the corresponding evolutions of
400 MAE, RMSE and NSE for both wind components, comparing the YRD1km product (red lines)
401 with the ERA5 reanalysis (blue lines), based on validation against ground-based observational data.

402 The YRD1km dataset consistently outperforms ERA5 across both components and all
403 metrics. MAE values for YRD1km remain consistently lower than those of ERA5, particularly
404 during nighttime hours, in agreement with the statistical results summarized in Table 4, which
405 show MAE reductions of 21.61% for U10m and 26.04% for V10m. In addition, the RMSE values
406 for U10m and V10m are reduced by 18.30% and 22.63%, respectively. These results indicate the
407 effectiveness of combining multi-source nudging and high-resolution land use data in consistently
408 capturing subtle wind variations over time.



409

410 *Figure 6. Time series of model performance metrics for hourly 10-m wind components over the*
 411 *YRD region in June 2022. Panels (a), (b) and (c) show the MAE, RMSE and NSE, respectively, for*
 412 *the U10m. Panels (d), (e) and (f) show the corresponding MAE, RMSE and NSE metrics for the*

413 *V10m. All metrics are computed for each hour by spatially averaging errors over all available*
414 *surface stations across the YRD region. The red and blue lines represent the YRD1km and ERA5*
415 *datasets, respectively.*

416 Both wind components exhibit pronounced diurnal cycles in MAE and RMSE, characterized
417 by peak errors during daytime, particularly around local noon, and reduced errors during nighttime.
418 This pattern reflects the influence of boundary layer dynamics, where daytime convective mixing
419 enhances wind variability and poses greater challenges for model accuracy, whereas nocturnal
420 stability leads to more predictable near-surface wind behavior. The persistence and regularity of
421 this fluctuation across the month highlight the necessity of capturing diurnal processes in high-
422 resolution simulations.

423 In terms of NSE, YRD1km maintains higher and more stable values throughout the month
424 for both U10m and V10m. Specifically, NSE values increase by 33.27% for U10m and 40.13% for
425 V10m compared to ERA5. While ERA5 frequently exhibits degraded performance, including
426 negative NSE values during high-variability periods, YRD1km often sustains NSE above 0.4, with
427 frequent peaks exceeding 0.6, especially during nocturnal hours. This reflects a markedly
428 improved temporal agreement between modeled and observed wind variations.

429 Overall, the consistent improvements observed across both horizontal wind components
430 confirm the robustness of the proposed downscaling framework. By effectively addressing both
431 synoptic-scale and diurnal-scale variability, the YRD1km dataset provides a substantially
432 enhanced representation of near-surface wind fields in a complex and highly urbanized region such
433 as the YRD.

434 *Table 4. Evaluation of 10-m wind field simulation performance over the YRD region in June*
435 *2022.*

Variable	Sample size	Indicator	Data		Improvement (%)
			ERA5	YRD1km	
U10m (m/s)	243280	MAE	1.333	1.045	21.61
		RMSE	1.766	1.443	18.30
		NSE	0.468	0.645	33.27
V10m (m/s)	243280	MAE	1.474	1.090	26.04
		RMSE	1.938	1.500	22.63
		NSE	0.407	0.645	40.13

436 4.3 Evaluation of Vertical Wind Profile Accuracy Using Radiosonde Observations

437 To comprehensively evaluate the vertical simulation performance of the YRD1km dataset,
438 radiosonde observations from the Baoshan station in Shanghai (ID: 58362) were used for the
439 month of June 2022 at 00 and 12 UTC. This station is situated at 31.39°N, 121.45°E with an
440 elevation of 3.3 m, within a densely built-up urban area. A comparative analysis was conducted
441 between YRD1km and ERA5 reanalysis data for wind speed accuracy within the 1000–100 hPa
442 pressure range, focusing on both Bias and RMSE metrics. The YRD1km dataset provides outputs
443 at 32 standard vertical levels, ranging from 1000 hPa near the surface to 10 hPa in the upper
444 atmosphere. These levels are obtained through vertical interpolation from the original 61 terrain-
445 following eta levels of the WRF model. Key pressure levels include: 1000, 975, 950, 925, 900,
446 875, 850, 825, 800, 775, 750, 700, 650, 600, 550, 500, 450, 400, 350, 300, 250, 225, 200, 175, 150,
447 125, 100, 70, 50, 30, 20, and 10 hPa.

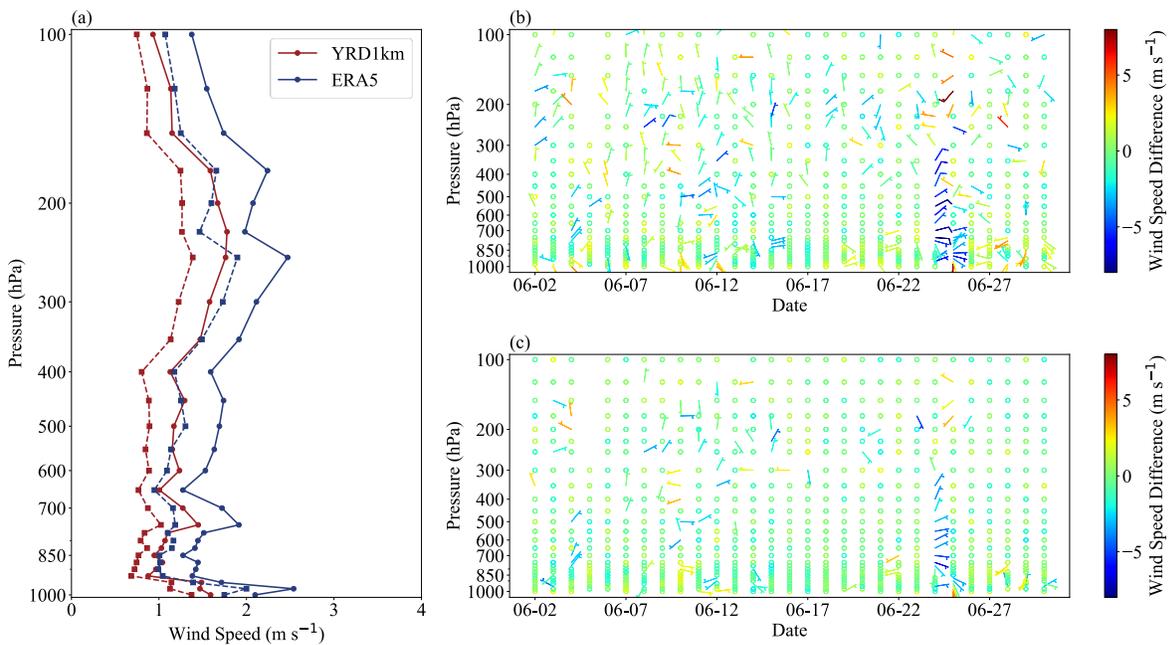
448 As illustrated in Figure 7a, the vertical profiles of MAE (dashed lines) and RMSE (solid
449 lines) reveal that the YRD1km dataset outperforms ERA5 across all pressure levels. The
450 improvements are pronounced in the lower troposphere, benefiting from the dynamic constraints
451 of multi-source observational nudging on near-surface winds and the refined land surface flux

452 representation driven by high-resolution land use data. The maximum reduction in RMSE reaches
453 up to 1.1 m/s at 975 hPa, representing a 42.18% improvement and highlighting the substantial
454 enhancement in near-surface wind speed accuracy provided by YRD1km.

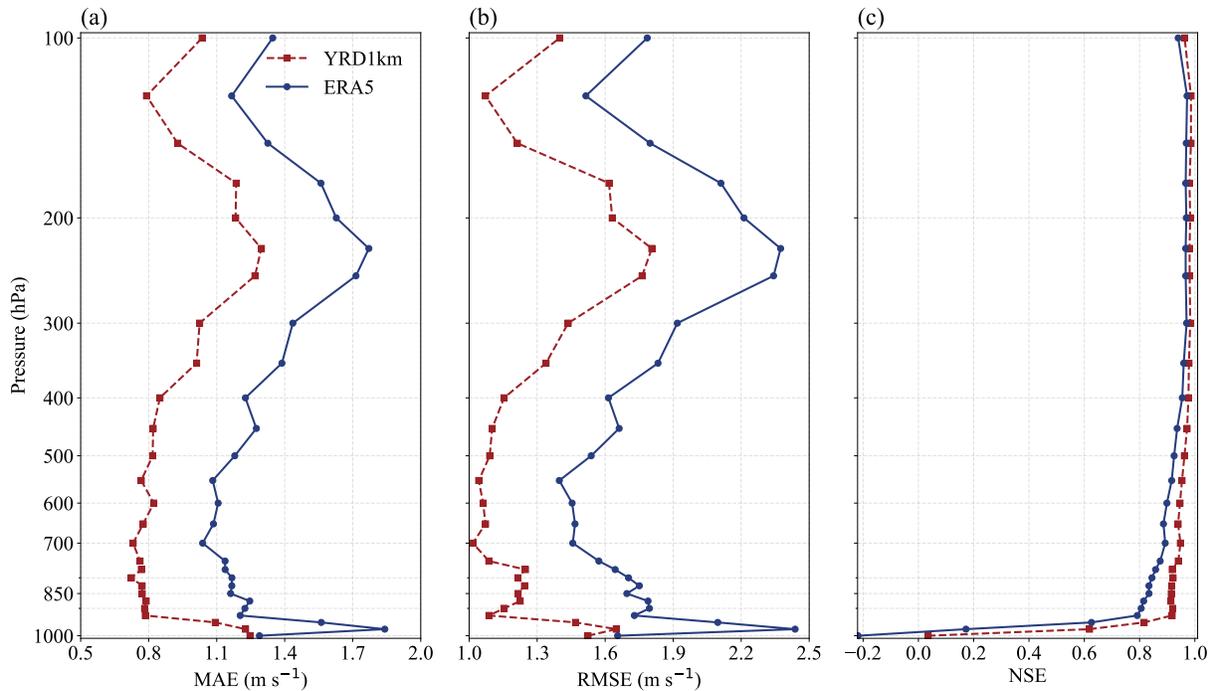
455 Time–height cross-section of wind vector differences plot (Figures 7b and 7c) further
456 highlights the clear performance of YRD1km. In Figure 7b, ERA5 exhibits frequent and large wind
457 speed differences, often exceeding ± 5 m/s, along with abrupt directional shifts, particularly within
458 the near-surface layer. Notably, at 00 UTC on June 24, radiosonde data indicate a sharp wind speed
459 increase above the 950 hPa level, exceeding 19.5 m/s, which ERA5 significantly underestimates.
460 This result is consistent with previous studies that have identified ERA5’s limitations in capturing
461 extreme wind events due to its coarser resolution and less-constrained boundary layer
462 parameterizations (Alkhalidi et al., 2025). In contrast, the YRD1km dataset exhibits a more stable
463 vertical wind structure, with smaller deviations from observed values. Although slight
464 underestimations remain during high wind episodes, the magnitude of extreme discrepancies is
465 considerably reduced compared to ERA5. This improvement underscores the effectiveness of the
466 multi-source observational nudging system in locally constraining vertical wind profiles and
467 enhancing model fidelity.

468 To further assess the representativeness of the vertical performance, we conducted similar
469 validations from all valid sounding samples at 11 radiosonde stations across the Yangtze River
470 Delta are now presented in Figure 8. These statistics show robust and systematic improvements of
471 YRD1km relative to ERA5, including average reductions of approximately 29.64% in MAE, 27.78%
472 in RMSE, and increases of about 47.75% in NSE, confirming that the enhanced 3D wind
473 performance of YRD1km is robust and regionally representative throughout the troposphere.

474 In summary, the YRD1km dataset, developed through the synergistic integration of high-
 475 resolution land surface information and multi-source data assimilation techniques, significantly
 476 improves not only near-surface wind simulations but also the representation of vertical wind
 477 structures. This provides a reliable, high-quality data foundation for a wide range of 3D wind field-
 478 dependent applications, such as low-level wind shear, wind turbine load estimation, pollutant
 479 cross-layer transport modeling, and urban atmospheric environment studies.



480
 481 *Figure 7. Vertical evaluation of wind field performance from the YRD1km and ERA5 datasets*
 482 *against radiosonde observations at the Baoshan station in Shanghai during June 2022. (a) Vertical*
 483 *profiles of wind speed MAE (dashed lines) and RMSE (solid lines) for YRD1km (red) and ERA5*
 484 *(blue), calculated from all available soundings at 00 and 12 UTC. (b) Time–height cross-section*
 485 *of wind vector differences between ERA5 and radiosonde observations (RAOB), with wind speed*
 486 *differences (m/s) indicated by color shading. (c) As in (b), but for YRD1km minus RAOB. Wind*
 487 *difference plots are shown at 24-hour intervals, beginning at 00 UTC on 2 June 2022.*



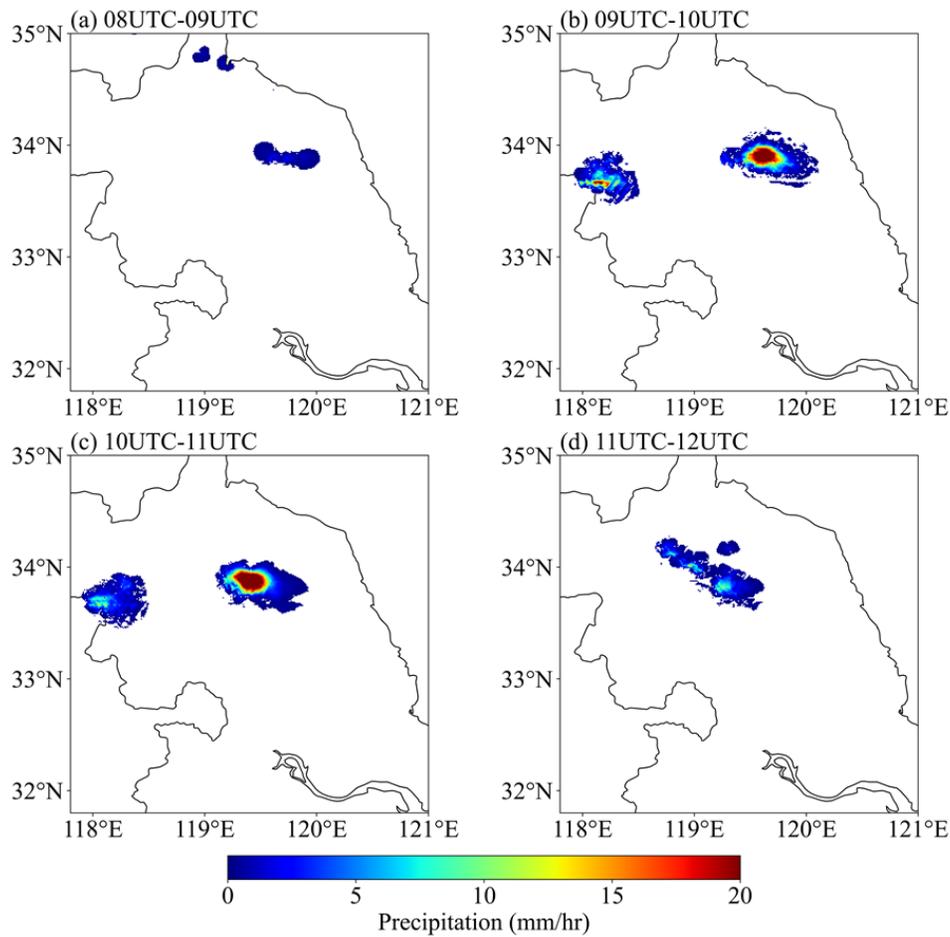
488

489 **Figure 8.** Vertical evaluation of wind field performance from the YRD1km and ERA5 datasets
 490 against radiosonde observations at 11 stations across the Yangtze River Delta in June 2022. Panels
 491 show (a) MAE, (b) RMSE and (c) NSE, computed using all available 00 UTC and 12 UTC
 492 soundings.

493 4.4 Case Studies of Severe Convection Events

494 While previous statistical validations have demonstrated the superior performance of the
 495 YRD1km dataset spatially and temporally, its advantages are further illustrated in short-term, high-
 496 impact convective events. In such situations, the high spatial and temporal resolution of YRD1km
 497 enables a more detailed diagnosis of mesoscale and vertical dynamical structures associated with
 498 convective initiation and evolution.

499 As illustrated in Figure 9, a convective storm outbreak occurred over northern Yancheng,
 500 Jiangsu Province, on the afternoon of 16 June 2022. The event was characterized by highly
 501 localized and intense precipitation, with peak hourly rainfall rates reaching up to $20 \text{ mm} \cdot \text{h}^{-1}$.



502

503 *Figure 9. Hourly evolution of precipitation associated with a convective storm over northern*

504

Yancheng, Jiangsu Province, on 16 June 2022.

505

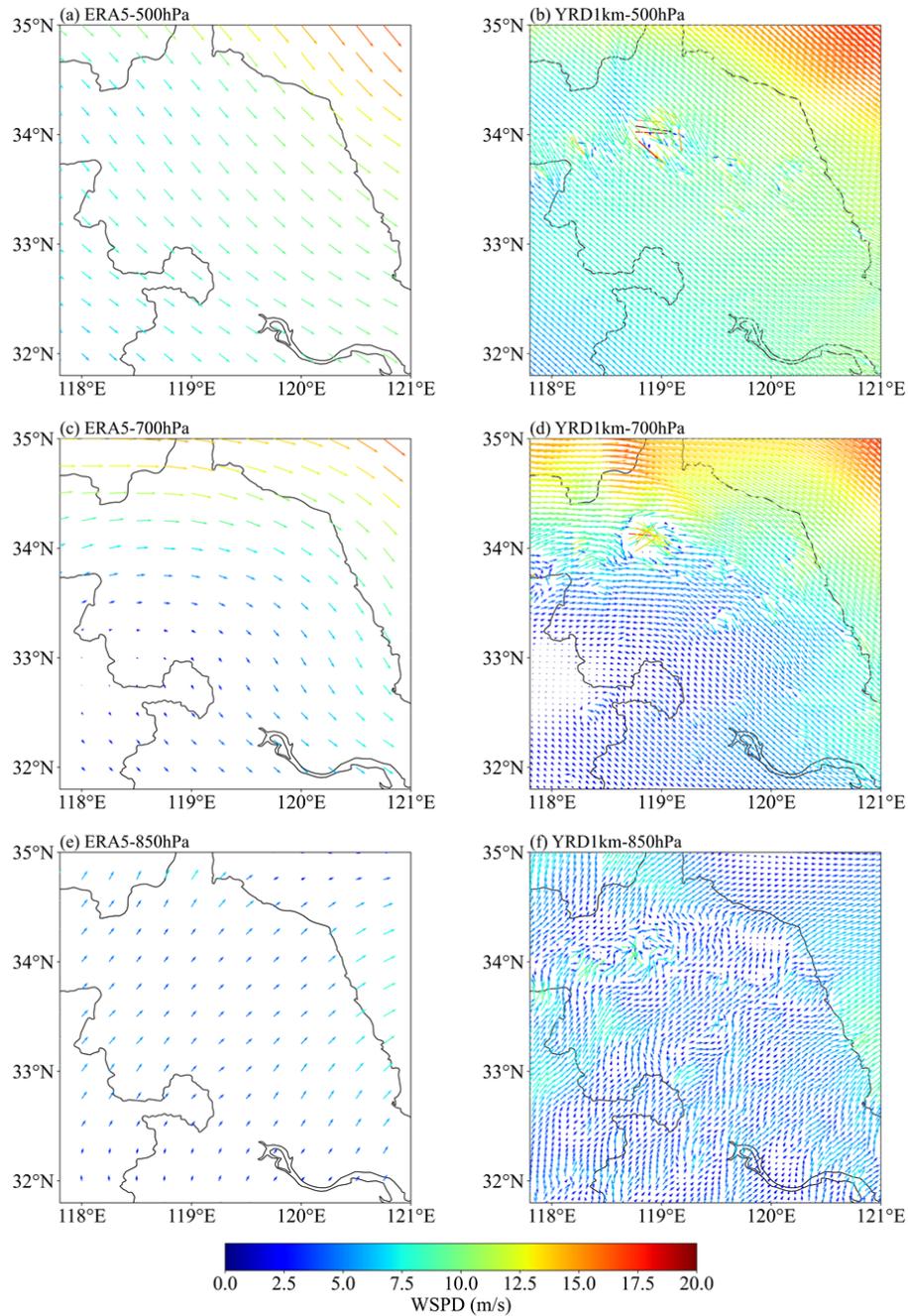
506

507

508

509

To investigate the applicability of the YRD1km dataset in high-impact weather scenarios, this study conducts a comparative analysis of wind field structures between ERA5 and YRD1km during the convective event, focusing on three key pressure levels: 500 hPa, 700 hPa, and 850 hPa (Figure 10). These levels are critical for identifying shear lines, low-level jets, and convective initiation mechanisms.

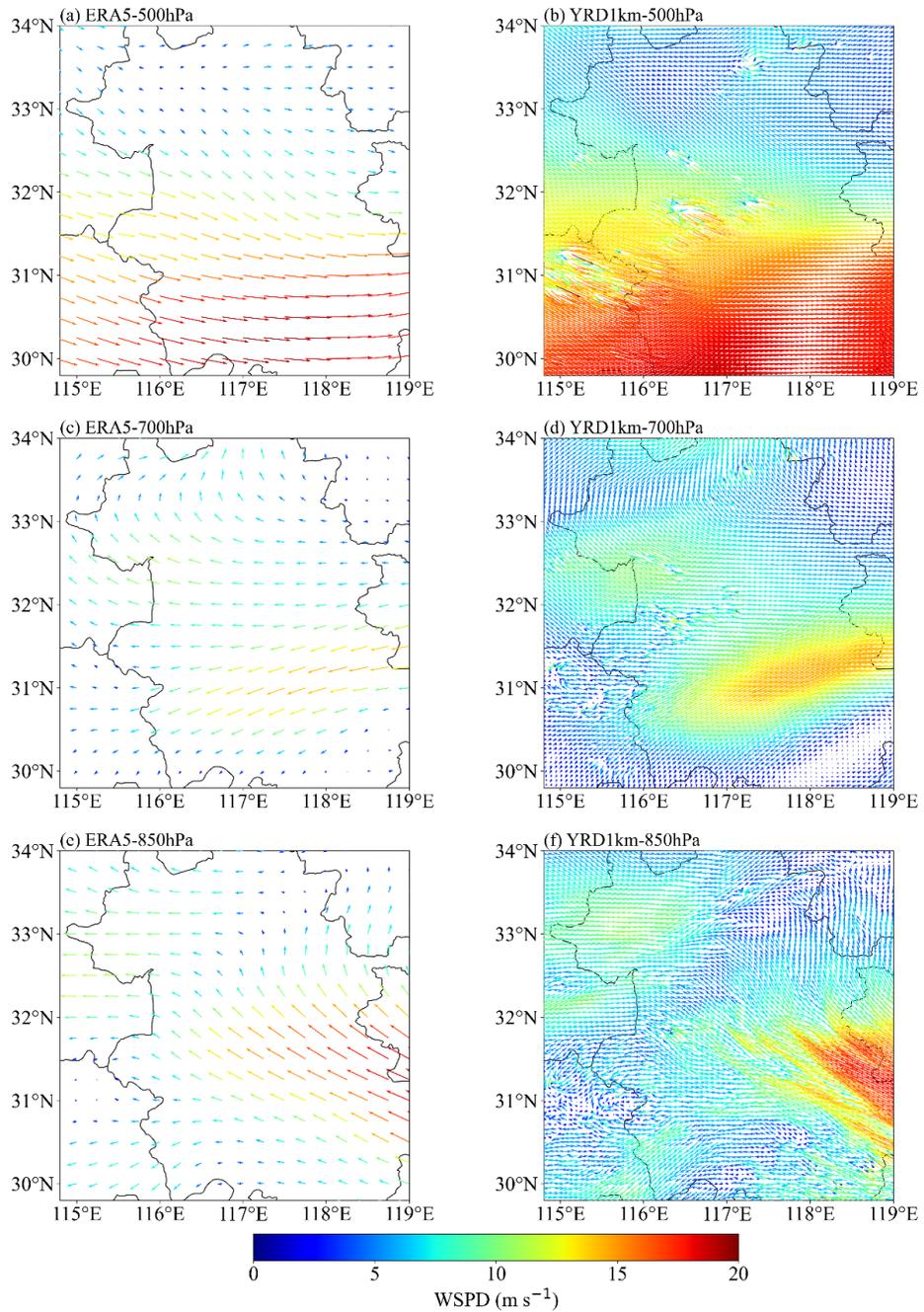


510

511 *Figure 10. Comparative analysis of wind field structures between the YRD1km and ERA5 datasets*
 512 *during the short-duration severe convective event over Yancheng, Jiangsu Province. Displayed are*
 513 *horizontal wind vectors (arrows) and wind speed (color shading) at the (a, b) 500 hPa, (c, d) 700*
 514 *hPa, and (e, f) 850 hPa levels from ERA5 (left column) and YRD1km (right column) at 08:00 UTC*
 515 *on 16 June 2022. For visual clarity, YRD1km wind vectors have been thinned by a factor of six.*

516 Overall, the wind field structure in ERA5 appears relatively homogeneous, limiting its ability
517 to capture mesoscale and sub-mesoscale disturbances. In contrast, YRD1km reveals more detailed
518 local structures and dynamic features, demonstrating a stronger capacity to resolve mesoscale
519 systems. Across all three pressure levels, YRD1km consistently captures regions of enhanced wind
520 speed, wind shear, and convergence. Notably, near 34°N, 119°E at 500 hPa, YRD1km identifies a
521 localized wind speed maximum exceeding 17.5 m/s and a well-defined shear zone. At 700 hPa, a
522 clear convergence band and wind speed enhancement area are observed, which is conducive to the
523 maintenance and development of the convective system. Although wind speeds weaken at 850 hPa,
524 perturbation signatures remain evident. These structural features spatially align with the center of
525 heavy precipitation during the event, indicating that YRD1km has enhanced diagnostic capability
526 in capturing the dynamical background for the initiation and maintenance of deep convective
527 systems.

528 Similar three-dimensional wind structures and dynamical consistency are also observed in an
529 independent mesoscale convective precipitation event on 10 June 2022 (Figure 11). During the
530 mature stage of this event, YRD1km consistently resolves enhanced wind speed gradients, low-
531 level convergence, and vertically coherent shear structures across the 500, 700, and 850 hPa levels,
532 whereas ERA5 depicts comparatively smooth and spatially homogeneous flow patterns. In
533 particular, YRD1km captures a well-defined low-level jet and its associated convergence zone at
534 850 hPa, together with mesoscale wind speed enhancements at mid-levels. The spatial
535 configuration of these features closely corresponds to the observed organization of convective
536 precipitation, highlighting the ability of YRD1km to represent dynamically relevant structures that
537 are critical for convective system development.



538

539 *Figure 11. Comparative analysis of horizontal wind field structures between ERA5 and YRD1km*

540 *during a mesoscale convective precipitation event over the study region on 10 June 2022. Shown*

541 *are horizontal wind vectors (arrows) and wind speed (color shading) at the (a, b) 500 hPa, (c, d)*

542 *700 hPa, and (e, f) 850 hPa levels from ERA5 (left column) and YRD1km (right column) at 04:00*

543 *UTC, corresponding to the mature stage of the convective system. For visual clarity, wind vectors*
544 *in YRD1km are thinned by a factor of six.*

545 Collectively, these case analyses demonstrate that the high spatial resolution of YRD1km
546 enables a physically consistent depiction of three-dimensional wind field structures during severe
547 convective events. Such capability supports its use in diagnostic studies of mesoscale and
548 convective-scale dynamics, and provides a valuable background wind field for applications
549 including high-resolution air-quality dispersion modeling, event-based wind hazard assessment,
550 and the development of training datasets for data-driven downscaling approaches.

551 **5. Conclusions**

552 This study developed and rigorously validated YRD1km, a high-resolution (1 km, hourly)
553 3D wind field dataset over the YRD region. The dataset was generated through dynamical
554 downscaling of ERA5 reanalysis data using a customized WRF model configuration. It was further
555 refined by integrating multi-source observational nudging and updated land use representations to
556 improve surface parameterization.

557 Comprehensive validations using surface station and radiosonde observations confirmed that
558 YRD1km significantly outperforms ERA5 in both near-surface and vertical wind simulations. For
559 10-m wind fields, YRD1km consistently achieved smaller errors and higher skill scores across
560 MAE, RMSE, and NSE, at both hourly and daily scales. The dataset also better characterizes
561 spatial variability in wind speed, particularly over complex terrain and densely urbanized areas.
562 Its wind vector fields align well with underlying geographic features, and monthly statistics show
563 reductions in MAE and RMSE of approximately 20%, with NSE improved by more than 33%. In
564 the vertical dimension, YRD1km exhibited reduced RMSE across nearly all pressure levels and
565 produced observation-consistent vertical profiles. *Analyses of severe convective events*

566 demonstrated YRD1km’s ability to resolve fine-scale dynamic signatures, including wind shear,
567 low-level convergence, and enhanced wind zones, supporting improved diagnosis of convective
568 development mechanisms.

569 These findings highlight the value of high-resolution wind datasets that integrate physically
570 consistent dynamical downscaling with observational constraints in representing mesoscale and
571 diurnal wind variability over complex urbanized environments. The YRD1km dataset provides a
572 robust, physically coherent foundation for applications that require detailed three-dimensional
573 wind information, including wind energy resource assessment, urban boundary-layer diagnostics,
574 and high-resolution air quality and pollutant dispersion modeling. In particular, the resolved fine-
575 scale vertical wind structures offer important potential for diagnosing low-level wind shear and
576 related dynamical features, which are critical for understanding urban atmospheric processes and
577 for supporting safety-oriented analyses relevant to low-altitude airspace operations in rapidly
578 developing metropolitan regions.

579 Looking ahead, the proposed framework can be further extended to generate longer-term,
580 seasonally continuous high-resolution wind datasets and applied to other regions with complex
581 terrain and heterogeneous land use. Future developments may also incorporate additional
582 observational constraints, such as satellite-based wind products and higher-frequency ground-
583 based remote sensing measurements. In addition, the YRD1km dataset can serve as a physically
584 consistent reference or training dataset for emerging data-driven downscaling approaches,
585 including machine learning models and sub-kilometer-scale simulations, thereby supporting next-
586 generation high-resolution atmospheric modeling and applied research.

587

588 **Data availability**

589 The YRD1km 3D wind field dataset is available at <https://doi.org/10.57760/sciencedb.23752>

590 (Zhang et al., 2025).

591

592 **Author contributions**

593 ZZ: data collection and processing; methodology and analysis; writing (original draft preparation).

594 YL: conceptualization; supervision; methodology and analysis; writing (original draft preparation,

595 review, and editing). XM, PX, and JZ: data collection. ZL, MM, DD, and BL: writing (review and

596 editing). JL: supervision; writing (review and editing).

597

598 **Competing interests**

599 The contact author has declared that none of the authors has any competing interests.

600

601 **Acknowledgments**

602 We gratefully acknowledge the ECNU Multifunctional Platform for Innovation 001 for providing

603 high-performance computing resources. We also thank the ECMWF for access to the ERA5

604 reanalysis data, the NCEP for the global upper air and surface weather observations, the CMA for

605 providing AWS data, and the ESA for the WorldCover 2020 land use data. Finally, we sincerely

606 thank the editor and anonymous reviewers for their insightful comments and constructive

607 suggestions, which greatly improved the quality of this manuscript.

608

609 **Financial support**

610 This research was supported by the Shanghai Science and Technology Program (Grant No.
611 U2542201 and 25ZR1401099) and the National Natural Science Foundation of China (Grant No.
612 U2142201).

613

614 **References**

615 Alkhalidi, M., Al-Dabbous, A., Al-Dabbous, S., and Alzaid, D.: Evaluating the Accuracy of the
616 ERA5 Model in Predicting Wind Speeds Across Coastal and Offshore Regions, *J. Mar. Sci. Eng.*,
617 13, 149, <https://doi.org/10.3390/jmse13010149>, 2025.

618 Anderson, J. T., Hardy, E. E., Roach, J. T., and Witmer, R. E.: A land use and land cover
619 classification system for use with remote sensor data, Professional Paper, U.S. Geological Survey,
620 <https://doi.org/10.3133/pp964>, 1976.

621 Bao, J., Feng, J., and Wang, Y.: Dynamical downscaling simulation and future projection of
622 precipitation over China, *J. Geophys. Res. Atmospheres*, 120, 8227–8243,
623 <https://doi.org/10.1002/2015JD023275>, 2015.

624 Bernini, L., Lagasio, M., Milelli, M., Oberto, E., Parodi, A., Hachinger, S., Kranzlmüller, D., and
625 Tartaglione, N.: Convection-permitting dynamical downscaling of ERA5 for Europe and the
626 Mediterranean basin, *Quarterly Journal of the Royal Meteorological Society*, 151, e5014,
627 <https://doi.org/10.1002/qj.5014>, 2025.

628 Boé, J., Terray, L., Habets, F., and Martin, E.: Statistical and dynamical downscaling of the Seine
629 basin climate for hydro-meteorological studies, *Int. J. Climatol.*, 27, 1643–1655,
630 <https://doi.org/10.1002/joc.1602>, 2007.

631 Collier, E. and Mölg, T.: BAYWRF: a high-resolution present-day climatological atmospheric
632 dataset for Bavaria, *Earth Syst. Sci. Data*, 12, 3097–3112, [https://doi.org/10.5194/essd-12-3097-](https://doi.org/10.5194/essd-12-3097-2020)
633 2020, 2020.

634 Daescu, D. N. and Langland, R. H.: Error covariance sensitivity and impact estimation with adjoint
635 4D-Var: theoretical aspects and first applications to NAVDAS-AR, *Q. J. R. Meteorol. Soc.*, 139,
636 226–241, <https://doi.org/10.1002/qj.1943>, 2013.

637 Dayon, G., Boé, J., and Martin, E.: Transferability in the future climate of a statistical downscaling
638 method for precipitation in France, *J. Geophys. Res. Atmospheres*, 120, 1023–1043,
639 <https://doi.org/10.1002/2014JD022236>, 2015.

640 De Bode, M., Hedde, T., Roubin, P., and Durand, P.: A Method to Improve Land Use
641 Representation for Weather Simulations Based on High-Resolution Data Sets—Application to

642 Corine Land Cover Data in the WRF Model, *Earth Space Sci.*, 10, e2021EA002123,
643 <https://doi.org/10.1029/2021EA002123>, 2023.

644 Dudhia, J.: Numerical Study of Convection Observed during the Winter Monsoon Experiment
645 Using a Mesoscale Two-Dimensional Model, *J. Atmospheric Sci.*, 46, 3077–3107,
646 [https://doi.org/10.1175/1520-0469\(1989\)046<3077:NSOCOD>2.0.CO;2](https://doi.org/10.1175/1520-0469(1989)046<3077:NSOCOD>2.0.CO;2), 1989.

647 Dujardin, J. and Lehning, M.: Wind-Topo: Downscaling near-surface wind fields to high-
648 resolution topography in highly complex terrain with deep learning, *Q. J. R. Meteorol. Soc.*, 148,
649 1368–1388, <https://doi.org/10.1002/qj.4265>, 2022.

650 Dupuy, F., Durand, P., and Hedde, T.: Downscaling of surface wind forecasts using convolutional
651 neural networks, *Nonlinear Process. Geophys.*, 30, 553–570, [https://doi.org/10.5194/npg-30-553-](https://doi.org/10.5194/npg-30-553-2023)
652 2023, 2023.

653 Fu, D., Liu, Y., Li, H., Liu, S., Li, B., Thapa, S., Yabo, S., Sun, X., Tang, B., Zuo, J., Qi, H., and
654 Tian, C.: Evaluating the Impacts of Land Cover and Soil Texture Changes on Simulated Surface
655 Wind and Temperature, *Earth Space Sci.*, 7, e2020EA001173,
656 <https://doi.org/10.1029/2020EA001173>, 2020.

657 Golzio, A., Ferrarese, S., Cassardo, C., Diolaiuti, G. A., and Pelfini, M.: Land-Use Improvements
658 in the Weather Research and Forecasting Model over Complex Mountainous Terrain and
659 Comparison of Different Grid Sizes, *Bound.-Layer Meteorol.*, 180, 319–351,
660 <https://doi.org/10.1007/s10546-021-00617-1>, 2021.

661 Gutowski, W. J., Ullrich, P. A., Hall, A., Leung, L. R., O'Brien, T. A., Patricola, C. M., Arritt, R.
662 W., Bukovsky, M. S., Calvin, K. V., Feng, Z., Jones, A. D., Kooperman, G. J., Monier, E., Pritchard,
663 M. S., Pryor, S. C., Qian, Y., Rhoades, A. M., Roberts, A. F., Sakaguchi, K., Urban, N., and
664 Zarzycki, C.: The Ongoing Need for High-Resolution Regional Climate Models: Process
665 Understanding and Stakeholder Information, *Bull. Am. Meteorol. Soc.*, 101, E664–E683,
666 <https://doi.org/10.1175/bams-d-19-0113.1>, 2020.

667 Harkey, M. and Holloway, T.: Constrained dynamical downscaling for assessment of climate
668 impacts, *J. Geophys. Res. Atmospheres*, 118, 2136–2148, <https://doi.org/10.1002/jgrd.50223>,
669 2013.

670 Hersbach, H., Bell, B., Berrisford, P., Hirahara, S., Horányi, A., Muñoz-Sabater, J., Nicolas, J.,
671 Peubey, C., Radu, R., Schepers, D., Simmons, A., Soci, C., Abdalla, S., Abellan, X., Balsamo, G.,
672 Bechtold, P., Biavati, G., Bidlot, J., Bonavita, M., De Chiara, G., Dahlgren, P., Dee, D.,
673 Diamantakis, M., Dragani, R., Flemming, J., Forbes, R., Fuentes, M., Geer, A., Haimberger, L.,
674 Healy, S., Hogan, R. J., Hólm, E., Janisková, M., Keeley, S., Laloyaux, P., Lopez, P., Lupu, C.,
675 Radnoti, G., de Rosnay, P., Rozum, I., Vamborg, F., Villaume, S., and Thépaut, J.-N.: The ERA5
676 global reanalysis, *Q. J. R. Meteorol. Soc.*, 146, 1999–2049, <https://doi.org/10.1002/qj.3803>, 2020.

677 Höhle, K., Kern, M., Hewson, T., and Westermann, R.: A comparative study of convolutional
678 neural network models for wind field downscaling, *Meteorol. Appl.*, 27, e1961,
679 <https://doi.org/10.1002/met.1961>, 2020.

680 Hoke, J. E. and Anthes, R. A.: The Initialization of Numerical Models by a Dynamic-Initialization
681 Technique, *Mon. Weather Rev.*, 104, 1551–1556, [https://doi.org/10.1175/1520-0493\(1976\)104<1551:tionmb>2.0.co;2](https://doi.org/10.1175/1520-0493(1976)104<1551:tionmb>2.0.co;2), 1976.

683 Hong, S.-Y., Noh, Y., and Dudhia, J.: A New Vertical Diffusion Package with an Explicit Treatment
684 of Entrainment Processes, *Mon. Weather Rev.*, 134, 2318–2341,
685 <https://doi.org/10.1175/mwr3199.1>, 2006.

686 Horvath, K., Koracin, D., Vellore, R., Jiang, J., and Belu, R.: Sub-kilometer dynamical
687 downscaling of near-surface winds in complex terrain using WRF and MM5 mesoscale models, *J.*
688 *Geophys. Res. Atmospheres*, 117, <https://doi.org/10.1029/2012JD017432>, 2012.

689 Hu, W., Scholz, Y., Yeligeti, M., Bremen, L. von, and Deng, Y.: Downscaling ERA5 wind speed
690 data: a machine learning approach considering topographic influences, *Environ. Res. Lett.*, 18,
691 094007, <https://doi.org/10.1088/1748-9326/aceb0a>, 2023.

692 Jung, C. and Schindler, D.: On the influence of wind speed model resolution on the global technical
693 wind energy potential, *Renew. Sustain. Energy Rev.*, 156, 112001,
694 <https://doi.org/10.1016/j.rser.2021.112001>, 2022.

695 Kain, J. S.: The Kain–Fritsch Convective Parameterization: An Update, *J. Appl. Meteorol.*, 43,
696 170–181, [https://doi.org/10.1175/1520-0450\(2004\)043<0170:TKCPAU>2.0.CO;2](https://doi.org/10.1175/1520-0450(2004)043<0170:TKCPAU>2.0.CO;2), 2004.

697 Kotlarski, S., Szabó, P., Herrera, S., Rätty, O., Keuler, K., Soares, P. M., Cardoso, R. M., Bosshard,
698 T., Pagé, C., Boberg, F., Gutiérrez, J. M., Isotta, F. A., Jaczewski, A., Kreienkamp, F., Liniger, M.
699 A., Lussana, C., and Pianko-Kluczyńska, K.: Observational uncertainty and regional climate
700 model evaluation: A pan-European perspective, *Int. J. Climatol.*, 39, 3730–3749,
701 <https://doi.org/10.1002/joc.5249>, 2019.

702 Lei, L. and Hacker, J. P.: Nudging, Ensemble, and Nudging Ensembles for Data Assimilation in
703 the Presence of Model Error, *Mon. Weather Rev.*, 143, 2600–2610, <https://doi.org/10.1175/MWR-D-14-00295.1>, 2015.

705 Lian, J., Huang, S., Shao, J., Chen, P., Tang, S., Lu, Y., and Yu, H.: TerraWind: A Deep Learning-
706 Based Near-Surface Winds Downscaling Model for Complex Terrain Region, *Geophys. Res. Lett.*,
707 51, e2024GL112124, <https://doi.org/10.1029/2024GL112124>, 2024.

708 Liu, J., Shi, C., Ge, L., Tie, R., Chen, X., Zhou, T., Gu, X., and Shen, Z.: Enhanced Wind Field
709 Spatial Downscaling Method Using UNET Architecture and Dual Cross-Attention Mechanism,
710 *Remote Sens.*, 16, 1867, <https://doi.org/10.3390/rs16111867>, 2024a.

711 Liu, S., Zeman, C., and Schär, C.: Dynamical Downscaling of Climate Simulations in the Tropics,
712 *Geophys. Res. Lett.*, 51, e2023GL105733, <https://doi.org/10.1029/2023GL105733>, 2024b.

713 Liu, Y., Di, D., Li, J., Li, Z., Ma, Z., Zheng, J., Liu, Y.-A., and Zhang, T.: Consistency Assessment
714 of the Winds in Reanalysis Datasets and the GIIRS Product Using Radiosondes, *Adv. Atmospheric*
715 *Sci.*, <https://doi.org/10.1007/s00376-025-4138-x>, 2025.

716 Lo, J. C., Yang, Z., and Pielke, R. A.: Assessment of three dynamical climate downscaling methods
717 using the Weather Research and Forecasting (WRF) model, *J. Geophys. Res. Atmospheres*, 113,
718 2007JD009216, <https://doi.org/10.1029/2007JD009216>, 2008.

719 Mlawer, E. J., Taubman, S. J., Brown, P. D., Iacono, M. J., and Clough, S. A.: Radiative transfer
720 for inhomogeneous atmospheres: RRTM, a validated correlated-k model for the longwave, *J.*
721 *Geophys. Res. Atmospheres*, 102, 16663–16682, <https://doi.org/10.1029/97JD00237>, 1997.

722 Molina, M. O., Gutiérrez, C., and Sánchez, E.: Comparison of ERA5 surface wind speed
723 climatologies over Europe with observations from the HadISD dataset, *Int. J. Climatol.*, 41, 4864–
724 4878, <https://doi.org/10.1002/joc.7103>, 2021.

725 Nash, J. E. and Sutcliffe, J. V.: River flow forecasting through conceptual models part I — A
726 discussion of principles, *J. Hydrol.*, 10, 282–290, [https://doi.org/10.1016/0022-1694\(70\)90255-6](https://doi.org/10.1016/0022-1694(70)90255-6),
727 1970.

728 Nashwan, M. S., Shahid, S., and Chung, E.-S.: Development of high-resolution daily gridded
729 temperature datasets for the central north region of Egypt, *Sci. Data*, 6, 138,
730 <https://doi.org/10.1038/s41597-019-0144-0>, 2019.

731 Sahu, R. K., Bangalath, H. K., Mostamandi, S., Evans, J., Kucera, P. A., and Beck, H. E.:
732 Evaluating microphysics and boundary layer schemes in WRF: assessment of 36 scheme
733 combinations for 17 major storms in Saudi Arabia, *EGU sphere* [preprint],
734 <https://doi.org/10.5194/egusphere-2025-912>, 2025.

735 Santos-Alamillos, F. J., Pozo-Vázquez, D., Ruiz-Arias, J. A., and Tovar-Pescador, J.: Influence of
736 land-use misrepresentation on the accuracy of WRF wind estimates: Evaluation of GLCC and
737 CORINE land-use maps in southern Spain, *Atmospheric Res.*, 157, 17–28,
738 <https://doi.org/10.1016/j.atmosres.2015.01.006>, 2015.

739 Siewert, J. and Kroszczynski, K.: Evaluation of High-Resolution Land Cover Geographical Data
740 for the WRF Model Simulations, *Remote Sens.*, 15, 2389, <https://doi.org/10.3390/rs15092389>,
741 2023.

742 Skamarock, W. C., Klemp, J. B., Dudhia, J., Gill, D. O., Liu, Z., Berner, J., Wang, W., Powers, J.
743 G., Duda, M. G., Barker, D. M., and Huang, X.-Y.: A Description of the Advanced Research WRF
744 Model Version 4, <https://doi.org/10.5065/1DFH-6P97>, 2019.

745 Sun, Y., Deng, K., Ren, K., Liu, J., Deng, C., and Jin, Y.: Deep learning in statistical downscaling
746 for deriving high spatial resolution gridded meteorological data: A systematic review, *ISPRS J.*
747 *Photogramm. Remote Sens.*, 208, 14–38, <https://doi.org/10.1016/j.isprsjprs.2023.12.011>, 2024.

748 Tang, J., Niu, X., Wang, S., Gao, H., Wang, X., and Wu, J.: Statistical downscaling and dynamical
749 downscaling of regional climate in China: Present climate evaluations and future climate
750 projections, *J. Geophys. Res. Atmospheres*, 121, 2110–2129,
751 <https://doi.org/10.1002/2015JD023977>, 2016.

752 Tareghian, R. and Rasmussen, P. F.: Statistical downscaling of precipitation using quantile
753 regression, *J. Hydrol.*, 487, 122–135, <https://doi.org/10.1016/j.jhydrol.2013.02.029>, 2013.

754 Tewari A. M., Chen A. F., Wang A. W., Dudhia A. J., LeMone A. M. A., Mitchell A. K. E., Ek A.
755 M. B., Gayno A. G., Wegiel A. J. W., Cuenca A. R., and Society S. A. M.: Implementation and
756 verification of the unified Noah land-surface model in the WRF model, 20th Conference on
757 Weather Analysis and Forecasting/16th Conference on Numerical Weather Prediction, Seattle, WA:
758 American Meteorological Society, 2004.

759 Thompson, G., Field, P. R., Rasmussen, R. M., and Hall, W. D.: Explicit Forecasts of Winter
760 Precipitation Using an Improved Bulk Microphysics Scheme. Part II: Implementation of a New
761 Snow Parameterization, *Mon. Weather Rev.*, 136, 5095–5115,
762 <https://doi.org/10.1175/2008mwr2387.1>, 2008.

763 Viterbo, F., Sperati, S., Vitali, B., D’Amico, F., Cavalleri, F., Bonanno, R., and Lacavalla, M.:
764 MERIDA HRES: A new high-resolution reanalysis dataset for Italy, *Meteorol. Appl.*, 31, e70011,
765 <https://doi.org/10.1002/met.70011>, 2024.

766 Xu, Z., Han, Y., Tam, C.-Y., Yang, Z.-L., and Fu, C.: Bias-corrected CMIP6 global dataset for
767 dynamical downscaling of the historical and future climate (1979–2100), *Sci. Data*, 8, 293,
768 <https://doi.org/10.1038/s41597-021-01079-3>, 2021.

769 Yuan, H., Cheng, L., Pan, Y., Tan, Z., Liu, Q., and Jin, Z.: A multi-level parallel approach to
770 increase the computation efficiency of a global ocean temperature dataset reconstruction, *J.*
771 *Parallel Distrib. Comput.*, 192, 104938, <https://doi.org/10.1016/j.jpdc.2024.104938>, 2024.

772 Zamo, M., Bel, L., Mestre, O., and Stein, J.: Improved Gridded Wind Speed Forecasts by Statistical
773 Postprocessing of Numerical Models with Block Regression, *Weather Forecast.*, 31, 1929–1945,
774 <https://doi.org/10.1175/waf-d-16-0052.1>, 2016.

775 Zhang, J., Liu, B., Ren, S., Han, W., Ding, Y., and Peng, S.: A 4 km daily gridded meteorological
776 dataset for China from 2000 to 2020, *Sci. Data*, 11, 1230, [https://doi.org/10.1038/s41597-024-](https://doi.org/10.1038/s41597-024-04029-x)
777 [04029-x](https://doi.org/10.1038/s41597-024-04029-x), 2024.

778 Zhang, L., Xu, Y., Meng, C., Li, X., Liu, H., and Wang, C.: Comparison of Statistical and Dynamic
779 Downscaling Techniques in Generating High-Resolution Temperatures in China from CMIP5
780 GCMs, *J. Appl. Meteorol. Climatol.*, 59, 207–235, <https://doi.org/10.1175/jamc-d-19-0048.1>,
781 2020.

782 Zhang, N., Gao, Z., Wang, X., and Chen, Y.: Modeling the impact of urbanization on the local and
783 regional climate in Yangtze River Delta, China, *Theor. Appl. Climatol.*, 102, 331–342,
784 <https://doi.org/10.1007/s00704-010-0263-1>, 2010.

785 Zhang, S. and Li, X.: Future projections of offshore wind energy resources in China using CMIP6
786 simulations and a deep learning-based downscaling method, *Energy*, 217, 119321,
787 <https://doi.org/10.1016/j.energy.2020.119321>, 2021.

788 Zhang, Z. and Liu, Y.-A.: A 1 km Hourly High-Resolution 3D Wind Field Dataset over the Yangtze
789 River Delta during June-August 2021-2023, Science Data Bank [data set],
790 <https://doi.org/10.57760/sciencedb.23752>, 2025.

791