

# First High-Resolution Surface Spectral Clear-Sky Ultraviolet Radiation Dataset across China (1981–2023): Development, Validation, and Variability

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**Abstract.** Solar ultraviolet radiation (UV) plays a fundamental role in the Earth's energy balance, influencing a wide range of processes, including material degradation, biophysical reactions, ecological dynamics, or public health. In this context, the first high-resolution ( $10\times 10$  km) hourly dataset of surface solar UV under clear-sky conditions over mainland China from 1981 to 2023 is introduced, derived from ERA5 and MERRA2 reanalysis data and a reconstruction based on the SMARTS (Simple Model of the Atmospheric Radiative Transfer of Sunshine) spectral model. Leveraging the SMARTS model's accuracy and capabilities, this dataset provides UV data at 0.5 nm intervals between 280nm and 400nm, offering enhanced granularity for wavelength-specific analysis, thus filling a key gap in high-resolution hourly UV data for China. Validation of the UV dataset against ground observations at 37 stations of the Chinese Ecosystem Research Network (CERN) demonstrates strong performance, with a correlation coefficient (R), root mean square error (RMSE), and mean bias error (MBE) of 0.919, 5.07 W m<sup>-2</sup> and -0.07 W m<sup>-2</sup>, respectively. Compared with the [Clouds and the Earth's Radiant Energy System \(CERES\)](#) UV product, this dataset offers higher spatial and temporal resolution as well as higher accuracy in comparison with observations, thus enhancing data quality for a wide range of applications. The spatial and temporal distribution of clear-sky UV radiation exhibits distinct regional and seasonal variations, with higher values in the west and south, and lower values in the east and north. Over the past 43 years, the annual mean clear-sky broadband UV radiation averaged over China was 20.05 W m<sup>-2</sup>, showing a slightly increasing trend (+0.0237 W m<sup>-2</sup> yr<sup>-1</sup>). This dataset is now available at

<https://doi.org/10.6084/m9.figshare.28234298> , offering a valuable resource for addressing regional challenges related to UV radiation.

## 1 Introduction

Although ultraviolet (UV) radiation accounts for a small fraction ( $\approx 8\%$ ) of the total solar radiation (Gueymard, 2004), it significantly impacts human health, ecosystems, and various atmospheric photochemical reactions (Neale et al., 2023; Thomas et al., 2012), in particular. A moderate level of UV radiation facilitates the synthesis of vitamin D3 in the human body, essential for calcium absorption and bones health (Novotná et al., 2024; Wu et al., 2025). However, prolonged exposure to UV radiation can be harmful to the skin and eyes (Ma et al., 2023; Narayanan et al., 2010). UV radiation not only directly damages surface cells of plant leaves but also penetrates aquatic environment, affecting phytoplankton and weakening photosynthesis, which impacts plant growth (Williamson et al., 2014). UV radiation in the troposphere can accelerate photochemical reactions in the near-surface layer, contributing to the formation of secondary pollutants (Goti et al., 2024). Consequently, acquiring accurate UV data is essential in various fields, including public health, climate modeling, photobiology, ecosystem monitoring, and environmental management.

In the 1980s, the American Antarctic Program and the National Science Foundation established a high-latitude UV monitoring and observation network (Booth et al., 1994). In parallel, the Australian Radiation Laboratory (ARL) monitored UV radiation with spectroradiometers (SRM) and broadband detectors at 18 sites in Australia and Antarctica (Roy et al., 1998). In the 1990s, China set up its UV observation network with the establishment of a Brewer UV SRM at both the Zhongshan Research Station in Antarctica and the Waliguan Global Atmosphere Station on the Qinghai-Tibet Plateau, which marked the beginning of the UV observation network in China (Bo et al., 2009). The Chinese Ecosystem Research Network (CERN) established in 2004, continues to provide high-quality long-term UV data (Tong et al., 2023). Despite these efforts, the number of UV monitoring stations remains sparse compared to conventional meteorological or radiometric stations, leaving many regions of China uncovered (Lu et al., 2023). To address this gap, various methods for UV estimation have been developed, primarily classified into station-based and satellite-based approaches (Tang et al., 2022).

Station-based methods estimate the UV irradiance through an empirical model that ingests local

meteorological and/or radiometric data (Laiwarin et al., 2023; Qin et al., 2020b). For instance, the  
60 normalized global surface irradiance (usually referred to as “clearness index”), solar zenith angle, and  
total ozone column data can be used to estimate UV radiation (Antón et al., 2011; Wang et al., 2013).  
These methods have been applied successfully in regions such as the Tibetan Plateau, the Pearl River  
Delta and the North China Plain in China (Xia et al., 2008; Gong et al., 2014; Peng et al., 2015). Another  
alternative method is to utilize the locally measured global solar irradiance and other meteorological  
65 variables as inputs to empirical models (Barbero et al., 2006; Habte et al., 2019). Whereas station-based  
methods are expected to provide accurate UV estimates near the station under scrutiny, their applicability  
is limited when attempting to generalize results to other regions. Furthermore, these methods are sensitive  
to local weather conditions and geographic variations, making them unfit to represent broader regional  
patterns (Qi et al., 2024). One solution is to use gridded input data derived from reanalysis models or  
70 from satellite observations. In particular, satellite-based methods (Jesus et al., 2023) can map large-scale,  
spatially continuous UV radiation based on observations from instruments like the Total Ozone Mapping  
Spectrometer (TOMS) (Chubarova et al., 2020; Zerefos et al., 2023), the Global Ozone Monitoring  
Experiment (GOME) (Kujanpää and Kalakoski, 2015; Parisi et al., 2021), the Ozone Monitoring  
Instrument (OMI) (Valappil et al., 2024; Zhang et al., 2019), the TROPOspheric Monitoring Instrument  
75 (TROPOMI) (Lakkala et al., 2020; Lamy et al., 2021), or the Fengyun-4 (FY-4) (Qin et al., 2023; Wang  
et al., 2024). Several algorithms have been developed to estimate UV from satellite data, including  
statistical methods (Katsambas et al., 1997; Laguarda and Abal, 2019; Pei and He, 2019), look-up table  
(LUT) methods (Leng et al., 2023; Su et al., 2005; Verdebout, 2000), and radiative transfer models (Janjai  
et al., 2010).

80 Radiative transfer models explicitly consider various atmospheric processes such as scattering and  
absorption from ozone, air molecules, clouds, or aerosols (Huang et al., 2019). For instance, Meerkötter  
et al. (1997) and Li et al. (2000) calculate UV using models based on the matrix-operator theory and the  
discrete ordinate method (DISORT), respectively. Both methods account for multiple scattering and  
absorption processes. Such models have been proven effective for UV radiation estimation, but are  
85 computationally intensive and sensitive to various uncertainties, such as cloud fractional cover or adverse  
weather conditions, making them less suitable for large-scale UV estimation over extended spatial and  
temporal scales (Wu et al., 2022a, 2024).

In the past, many UV datasets containing long-term time series have been reconstructed based on satellite data (Ciren and Li, 2003; Čížková et al., 2018; Fragkos et al., 2024; Pei and He, 2019). Over  
90 China, Liu et al. (2017) reconstructed daily UV data from 1961 to 2014 using an all-sky estimation model combined with a hybrid model partly based on satellite observations. However, the main model's inputs use station-based meteorological data, resulting in an incomplete spatial coverage and a low spatiotemporal resolution. An attractive alternative consists in using gridded reanalysis data because of their global coverage, consistency, and perfect spatiotemporal continuity (no data breaks, contrary to  
95 satellite-based data). Reanalyses of interest include ERA5 (Li et al., 2023; Xia et al., 2021), MERRA2 (Laguarda et al., 2024; Lipponen et al., 2020), JRA-55 (Japanese 55-year Reanalysis) (Krizan, 2024; Wang et al., 2019) and CFSR (Climate Forecast System Reanalysis) (Li et al., 2024; Wang et al., 2011). Qin et al. (2020a) utilized MERRA2 to establish a novel physical broadband parameterization (FASTUV) for estimating surface solar UV radiation under all-sky conditions. Wu et al. (2022b) incorporated surface  
100 pressure and surface solar radiation from the ERA5 reanalysis data as input variables into their clear-sky UV radiation estimation model. Jiang et al. (2024) also used various ERA5 predictions, along with machine-learning methods, to derive the all-sky UV radiation over China on a daily basis.

Compared to the conventional methods reviewed above that only provide the *broadband* UV, the novelty of this study's goal is to obtain gridded estimates of the *spectral* UV irradiance under clear-sky  
105 conditions, and its spatiotemporal variations over mainland China. The remainder of this article is structured as follows. The methodology for estimating clear-sky UV radiation is presented in Section 2. Section 3 describes the input data used for that task, in addition to the observations and satellite data used for validation. Section 4 presents the validation results for the clear-sky UV radiation data, along with an analysis of the model's sensitivity and its spatial distribution over China. Section 5 discusses data  
110 availability, and finally Section 6 provides a summary and conclusions.

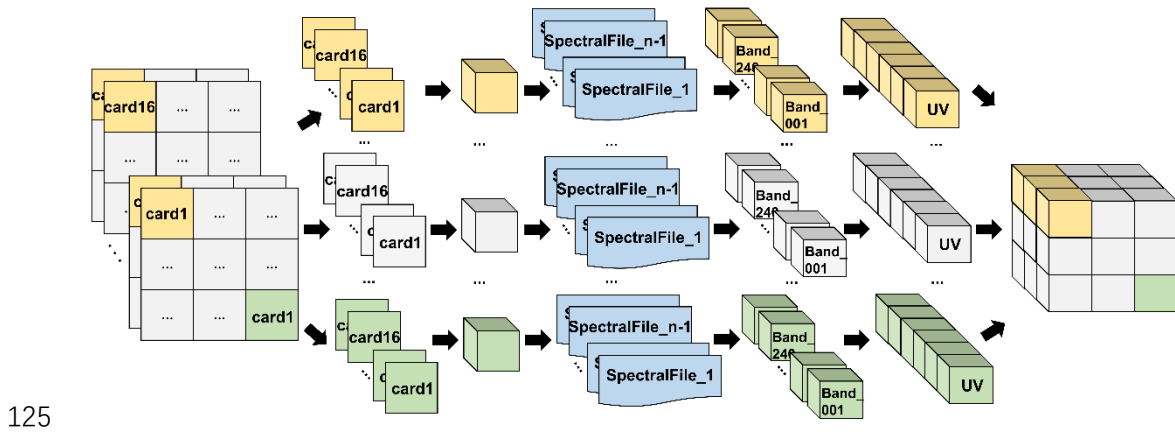
## 2 Methodology<sup>s</sup>

### 2.1 The SMARTS model

The “Simple Model for Atmospheric Transmission of Sunshine”, or SMARTS for short, is a radiative transfer spectral model that has been under development since the early 1990s (Gueymard, 2001,  
115 2005, 2019). SMARTS estimates the different clear-sky irradiance components (direct, diffuse, and

global) at the surface over the entire short-wave solar spectrum (280–4000 nm), and has become an essential tool in various disciplines, most importantly in various solar energy applications (Bicer et al., 2022; Mouhib et al., 2022; Pelland and Gueymard, 2022), the development of spectral irradiance standards (Habte et al., 2020; Xue and Igari, 2023), and UV research (Habte et al., 2019; N.Apell and McNeill, 2019). The UV band under scrutiny here extends from 280 to 400 nm at 0.5-nm resolution.

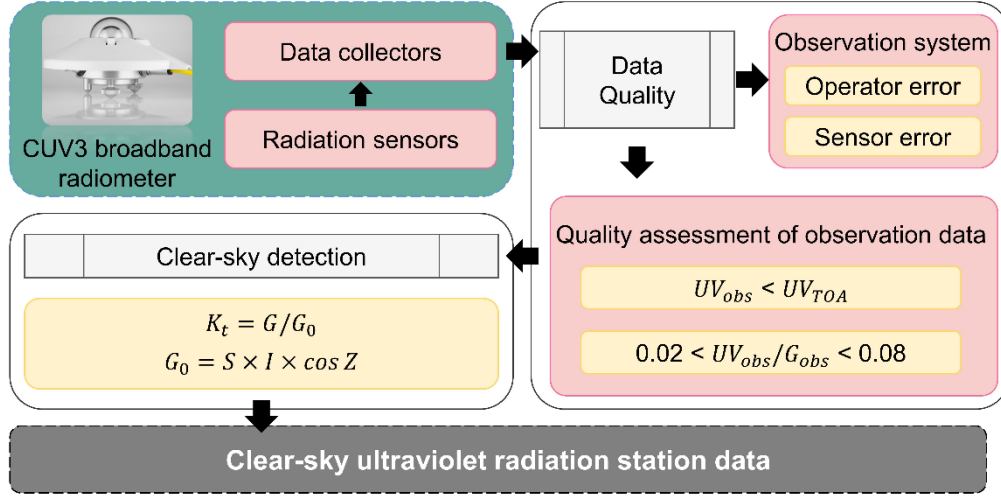
The SMARTS computational flow is shown in Figure 1. As visualized, single-image metadata is converted from planar to 3D at run time, thus significantly increasing the time cost of computation. To optimize the process, the model’s Fortran code was reconstructed and converted to MATLAB®. In that optimized implementation, the model can perform matrix operations in an efficient way.



**Figure 1. Schematic diagram of the SMARTS workflow process for the large-scale application involved here.**

## 2.2 Quality control of UV radiation observations

The quality control of the measured UV data involves two parallel procedures: (i) quality control of the observations; and (ii) detection of the clear-sky periods in the observational time series (Figure 2). The quality control of observation data can be divided into two parts: detection and elimination of systematic errors, and quality assessment of the observation data. The systematic errors includes operator errors and sensor errors. Furthermore, many obvious errors and missing observations need to be eliminated. The quality assessment of observational data is conducted in successive steps. First, the observed surface broadband UV irradiance should be less than extraterrestrial UV radiation. Second, the ratio of UV radiation to total solar radiation should be in the range 0.02–0.08 (Liu et al., 2017).



**Figure 2. Flowchart describing the quality control procedure of UV radiation observational data.**

It should be noted that, in the present investigation, the estimated UV radiation exclusively relates to clear-sky conditions. Therefore, it is necessary to conduct an efficient clear-sky screening of the observational time series, which needs to be done. The clearness index ( $K_t$ ), which is the ratio between the surface global solar irradiance ( $G$ ) measured at 1 hourly resolution and its extraterrestrial counterpart ( $G_0$ ), is employed as the primary screening tool in this study. Clear-sky conditions are identified as those for which  $K_t$  exceeds 0.7 (Qi et al., 2024). The global solar irradiance was obtained from the CERN site observations at 1 hourly temporal resolution and the extraterrestrial radiation was calculated using:

$$G_0 = S \times I \times \cos Z \quad (1)$$

where  $S$  is the sun-earth distance correction factor,  $I$  is the solar constant ( $1361.1 \text{ W m}^{-2}$ ) (Gueymard, 2018), and  $Z$  is the solar zenith angle. At each instant,  $Z$  and  $S$  are provided by a precise sun position algorithm.

### 2.3 Evaluation of the model

In the absence of spectral measurements, which would necessitate costly SRMs, the assessment of the accuracy of the clear-sky UV modeled estimates is done relatively to broadband measurements, as described in Section 3.1. The assessment uses conventional evaluation metrics (Gueymard, 2014), including mean bias error (MBE), root mean square error (RMSE), and correlation coefficient (R). MBE and RMSE are expressed both in absolute terms ( $\text{W m}^{-2}$ ) and relative percentage.

In addition, the probability density functions (PDFs) of the bias and the cumulative distribution functions (CDFs) of the absolute relative error (ARE) are used to compare the accuracy of the estimates

with respect to various satellite products (Jiang et al., 2020). The PDFs and CDFs provide complementary information about the ~~concentration and~~ distribution of the deviations between different UV radiation products and station observations. In particular, the position and shape of the PDF peaks can offer insights into the temporal characteristics of the bias. The CDF curves of different UV radiation products also allow a visual comparison of their overall performance. The steeper the CDF curve, the more concentrated the error and the higher the accuracy of the product, whereas a flat curve indicates a dispersed error and lower precision.

### 3 Data

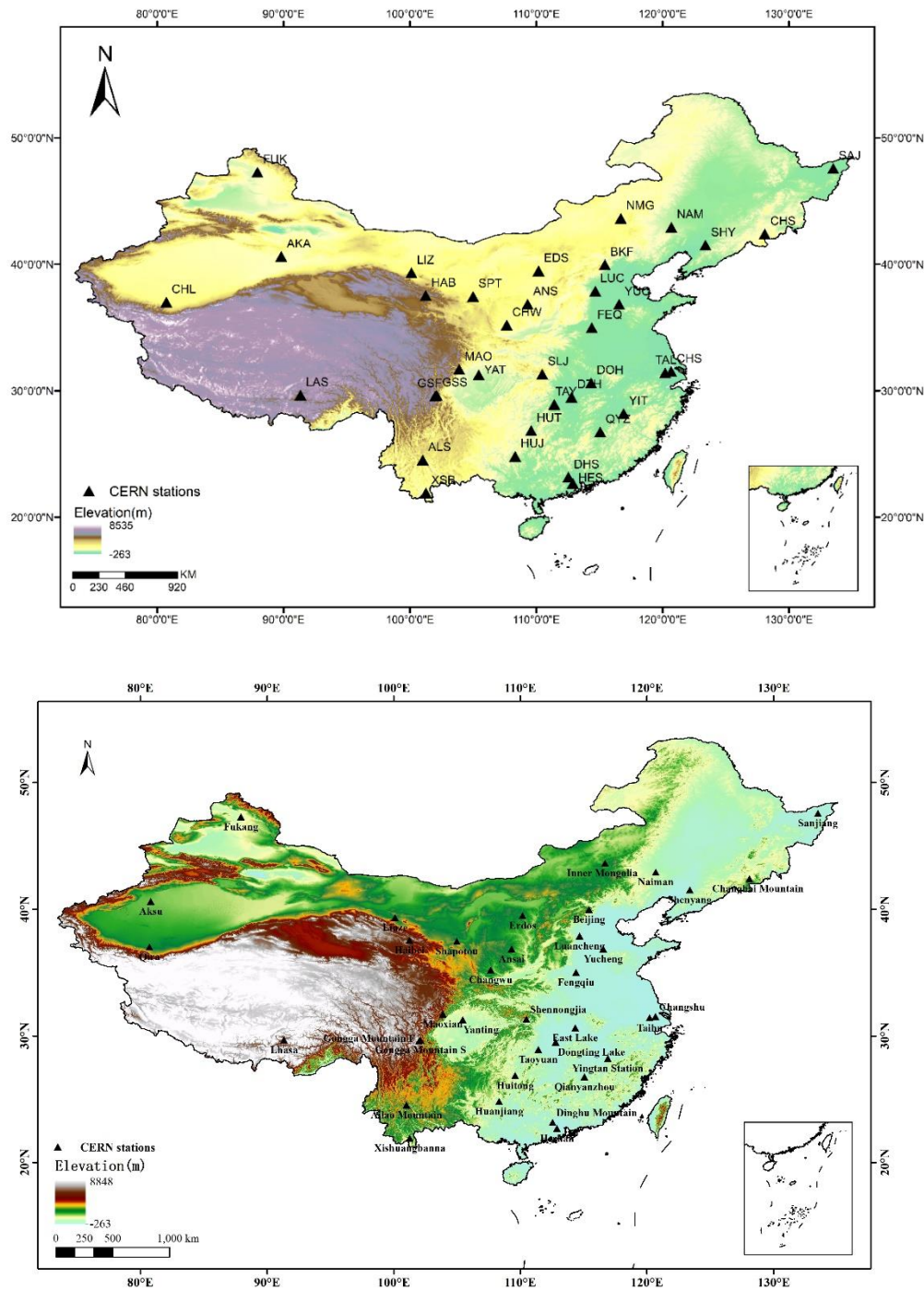
#### 3.1 Observational data from CERN

The UV database collected from the China Ecosystem Research Network (CERN) observation sites is used to validate the accuracy of the UV radiation estimates obtained with the SMARTS model. [Detailed information on CERN sites is provided in Table 1.](#) CERN has established a network of ecosystem observatories throughout China to monitor the meteorological, environmental, and radiometric conditions over a widely diverse range of ecosystems, including nature reserves, forest ecosystems, grassland ecosystems, wetland ecosystems, and urban ecosystems. The observational database selected for this study comprises hourly UV and total solar irradiance data from all the 37 stations in mainland China over the 2005-2013 period. The CERN stations use the CM11 global radiometer (Kipp & Zonen, The Netherlands) to measure global radiation with a spectral range of 285-2800 nm, with an accuracy of 5%. The CUV3 broadband radiometer (Kipp & Zonen, The Netherlands) is used to measure UV radiation, with a spectral range of 280-400 nm and an accuracy of 5%, which meets the World Meteorological Organization's (WMO) measurement standards. As shown in Figure 3, the CERN sites are relatively well distributed in the east. Due to the influence of topography, however, the density of sites in the west is much lower than in the east, which results in an uneven spatial distribution overall. [The spatial coverage of this product is constrained by the continuous availability of ERA5 reanalysis data, resulting in the exclusion of certain coastal observation stations at CERN.](#)

**Table 1. Detailed information about the CERN station used by this study.**

<u>Station</u>	<u>Latitude °N</u>	<u>Longitude °E</u>	<u>Climate Classification</u>
<u>Sanjiang</u>	<u>47.58</u>	<u>133.52</u>	<u>Temperate Monsoon Climate</u>
<u>Donghu</u>	<u>30.62</u>	<u>114.35</u>	<u>Subtropical Monsoon Climate</u>
<u>Linze</u>	<u>39.33</u>	<u>100.12</u>	<u>Temperate Continental Climate</u>
<u>Huitong</u>	<u>26.85</u>	<u>109.60</u>	<u>Subtropical Monsoon Climate</u>
<u>Neimenggu</u>	<u>43.63</u>	<u>116.70</u>	<u>Temperate Continental Climate</u>
<u>Beijing</u>	<u>39.98</u>	<u>115.43</u>	<u>Temperate Monsoon Climate</u>
<u>Qianyanzhou</u>	<u>26.75</u>	<u>115.07</u>	<u>Subtropical Monsoon Climate</u>
<u>Ailaoshan</u>	<u>24.53</u>	<u>101.02</u>	<u>Tropical Monsoon Climate</u>
<u>Taihu</u>	<u>31.42</u>	<u>120.22</u>	<u>Subtropical Monsoon Climate</u>
<u>Naiman</u>	<u>42.93</u>	<u>120.70</u>	<u>Temperate Continental Climate</u>
<u>Ansai</u>	<u>36.86</u>	<u>109.32</u>	<u>Temperate Monsoon Climate</u>
<u>Fengqiu</u>	<u>35.00</u>	<u>114.40</u>	<u>Temperate Monsoon Climate</u>
<u>Changshu</u>	<u>31.53</u>	<u>120.68</u>	<u>Subtropical Monsoon Climate</u>
<u>Lasa</u>	<u>29.67</u>	<u>91.33</u>	<u>plateau-highland climate</u>
<u>Luancheng</u>	<u>37.88</u>	<u>114.68</u>	<u>Temperate Monsoon Climate</u>
<u>Taoyuan</u>	<u>28.92</u>	<u>111.43</u>	<u>Subtropical Monsoon Climate</u>
<u>Shenyang</u>	<u>41.52</u>	<u>123.40</u>	<u>Temperate Monsoon Climate</u>
<u>Shapotou</u>	<u>37.47</u>	<u>105.00</u>	<u>Temperate Continental Climate</u>
<u>Dongtinghu</u>	<u>29.50</u>	<u>112.80</u>	<u>Subtropical Monsoon Climate</u>
<u>Haibei</u>	<u>37.53</u>	<u>101.25</u>	<u>plateau-highland climate</u>
<u>Huanjiang</u>	<u>24.82</u>	<u>108.33</u>	<u>Subtropical Monsoon Climate</u>
<u>Yanting</u>	<u>31.27</u>	<u>105.45</u>	<u>Subtropical Monsoon Climate</u>
<u>Shennongjia</u>	<u>31.32</u>	<u>110.48</u>	<u>Subtropical Monsoon Climate</u>
<u>Yucheng</u>	<u>36.85</u>	<u>116.57</u>	<u>Temperate Monsoon Climate</u>
<u>Cele</u>	<u>37.02</u>	<u>80.72</u>	<u>Temperate Continental Climate</u>
<u>Maoxian</u>	<u>31.70</u>	<u>103.90</u>	<u>Subtropical Monsoon Climate</u>
<u>Xishuangbanna</u>	<u>21.92</u>	<u>101.27</u>	<u>Tropical Monsoon Climate</u>
<u>GonggashanF</u>	<u>29.58</u>	<u>102.00</u>	<u>Subtropical Monsoon Climate</u>
<u>GonggashanS</u>	<u>29.65</u>	<u>102.12</u>	<u>Subtropical Monsoon Climate</u>
<u>Eerduosi</u>	<u>39.48</u>	<u>110.18</u>	<u>Temperate Continental Climate</u>
<u>Changwu</u>	<u>35.20</u>	<u>107.67</u>	<u>Temperate Monsoon Climate</u>
<u>Changbaishan</u>	<u>42.40</u>	<u>128.10</u>	<u>Temperate Monsoon Climate</u>
<u>Fukang</u>	<u>47.29</u>	<u>87.93</u>	<u>Temperate Continental Climate</u>
<u>Akesu</u>	<u>40.62</u>	<u>80.83</u>	<u>Temperate Continental Climate</u>
<u>Heshan</u>	<u>22.68</u>	<u>112.90</u>	<u>Tropical Monsoon Climate</u>
<u>Yingtian</u>	<u>28.20</u>	<u>116.92</u>	<u>Subtropical Monsoon Climate</u>
<u>Dinghushan</u>	<u>23.17</u>	<u>112.55</u>	<u>Tropical Monsoon Climate</u>





**Figure 3. Distribution of the 37 CERN sites in China that provide solar and UV irradiance observations.**

### 190 3.2 Reanalysis-derived SMARTS model input variables

The inputs to SMARTS are derived primarily from key products generated by the ERA5 and MERRA2 reanalysis over the period 1981–2023. ERA5 (ECMWF Reanalysis: Fifth generation) is an atmospheric reanalysis dataset published by the European Centre for Medium-Range Weather Forecasts

(ECMWF) to describe the state and variability of the Earth's atmospheric system (Muñoz-Sabater et al., 2021). The data employed in this study comprise ERA5\_Land, ERA5\_Pressure levels, and ERA5\_Single levels. ERA5\_Land has a spatial resolution of  $0.1^\circ \times 0.1^\circ$ , whereas ERA5\_Pressure levels and ERA5\_Single levels have a spatial resolution of  $0.25^\circ \times 0.25^\circ$ . Consequently, that coarser grid is resampled to achieve a unified spatial resolution of  $0.1^\circ \times 0.1^\circ$  or  $\approx 10$  km. The forecast albedo (FAL) from the ERA5-Land product is a land-specific parameter, as its values over ocean and lake surfaces are systematically set to NAN, making it suitable for continental-scale analysis.

MERRA-2 (Modern-Era Retrospective analysis for Research and Applications, Version 2) is an atmospheric reanalysis dataset released by NASA (Gelaro et al., 2017). MERRA-2 provides a long time series of AOD data (1980–present) with reasonable accuracy (Gueymard and Yang, 2020; Ou et al., 2022). The coarse resolution ( $0.5^\circ \times 0.625^\circ$ ) of the MERRA2 AOD data requires bilinear resampling to  $0.1^\circ \times 0.1^\circ$ , however. Table 42 shows the summarized information pertaining to all input variables. The position of each grid cell (latitude, longitude, and elevation) is also necessary to operate the radiation model.

**Table 42. Summarized information about the gridded variables used as inputs in this study.**

Dataset name	Parameters name	Resolution	
		Spatial	Temporal
ERA5-Pressure levels	Relative humidity (RH)	$0.25^\circ \times 0.25^\circ$	Hourly
ERA5-Single levels	Total column ozone (TCO3),	$0.25^\circ \times 0.25^\circ$	Hourly
	Total column water vapor (TCWV)		
ERA5-Land	Surface pressure (SP),	$0.1^\circ \times 0.1^\circ$	Hourly
	Forecast albedo (FAL),		
	2m temperature (T2M)		
MERRA-2	Aerosol optical depth (AOD)	$0.5^\circ \times 0.625^\circ$	Hourly

### 3.3 Satellite-based solar UV radiation product

CERES (Clouds and the Earth's Radiant Energy System) is a project initiated by NASA with the objective of conducting long-term monitoring and research on the Earth's radiant energy balance. In particular, CERES SYN1deg is a synthetic data base that provides various products related to the surface radiant field including radiative fluxes, cloudiness, and temperature. The SYN1deg database provides

global coverage at a horizontal resolution of  $1^\circ \times 1^\circ$ , with hourly temporal resolution (Wielicki et al.,  
215 1998). Here, the hourly UVA and UVB estimates from SYN1deg are used as a benchmark against which  
the present model's estimates can be assessed.

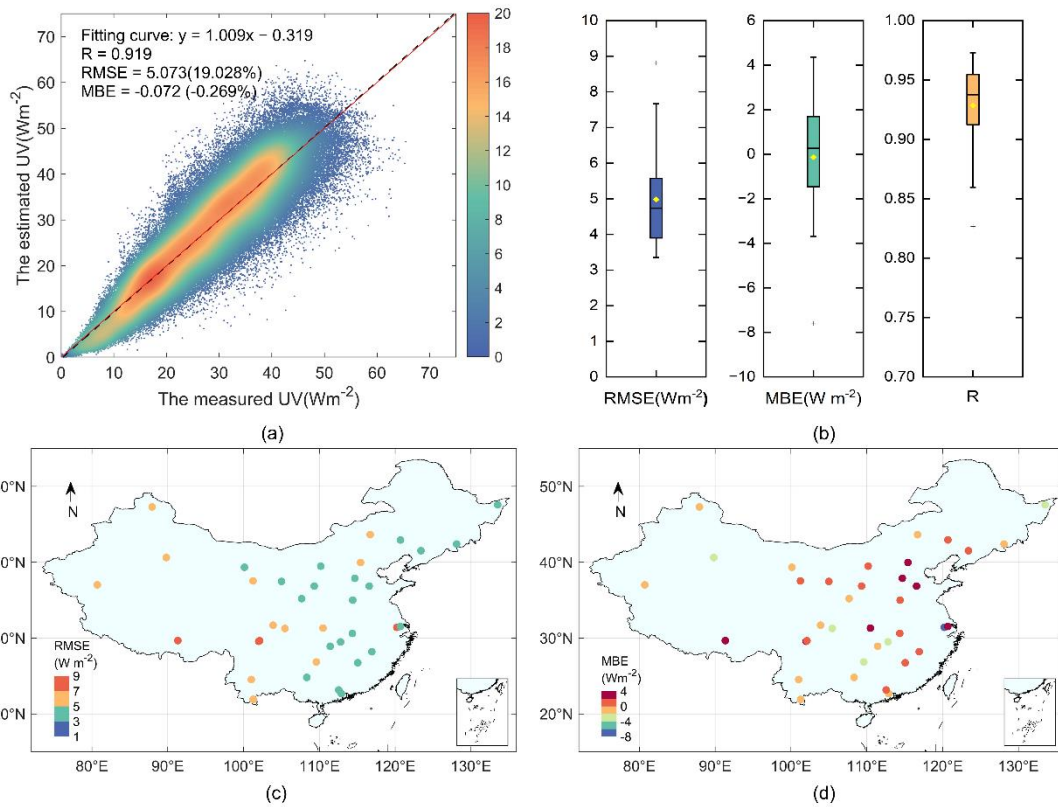
## 4 Results and discussion

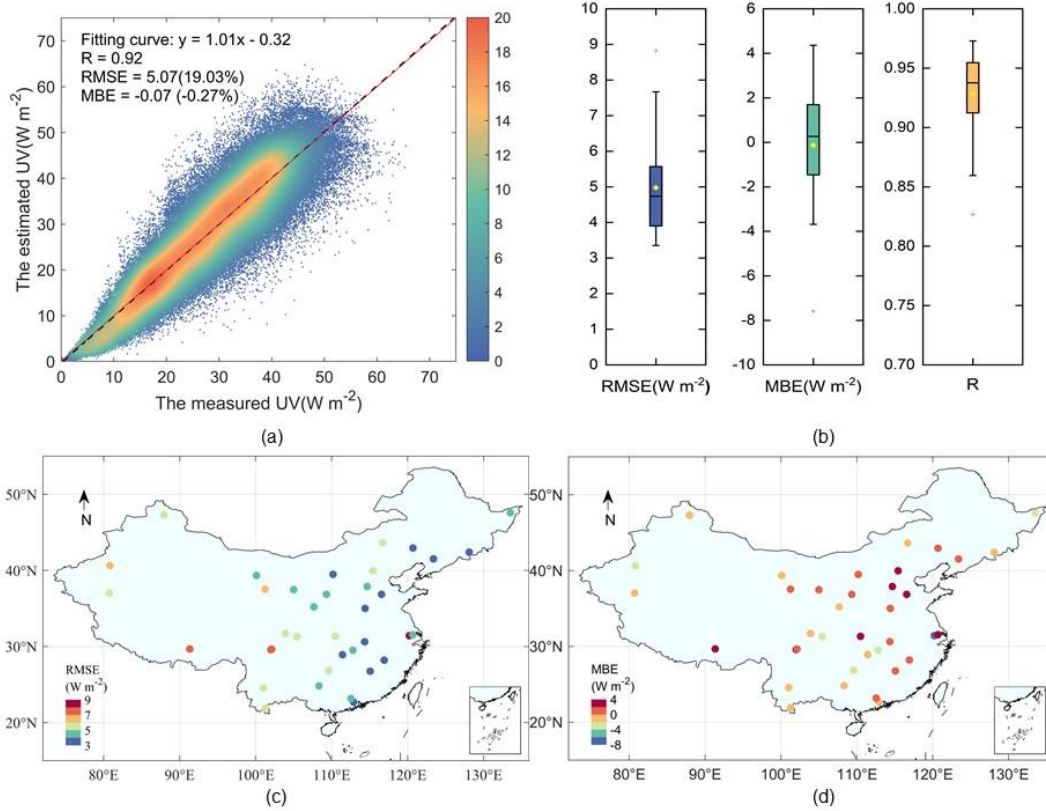
### 4.1 SMARTS-derived UV irradiance performance assessment

Figure 4 presents the validation results of the hourly clear-sky UV radiation model estimates,  
220 ~~compared with~~ Out of the 196,170 original observed data points from 37 CERN stations during 2005–  
2013, 170 were selected after applying the clear-sky screening procedure described in Section 2.2.  
Overall, the estimated hourly clear-sky broadband UV irradiance exhibits satisfactory accuracy, as  
evidenced by Figure 4a, with an R of 0.919, an RMSE of  $5.07 \text{ W m}^{-2}$ , and an MBE of  $-0.07 \text{ W m}^{-2}$ . The  
density scatter plot shows that the model performs well across most UV radiation levels, particularly in  
225 the range of  $20\text{--}40 \text{ W m}^{-2}$ , where the highest density of points is observed. However, at lower ( $<10 \text{ W m}^{-2}$ )  
and higher ( $>50 \text{ W m}^{-2}$ ) UV radiation levels, the scatter points become more dispersed, indicating  
increased model uncertainty under these extreme conditions. Notably, the regression slope of 1.009 and  
a negative intercept ( $-0.319$ ) suggest a slight systematic underestimation under low UV radiation  
conditions, possibly due to factors like local atmospheric characteristics not fully captured by the model.  
230 In addition, the upper and lower boundaries of the scatter points reveal distinct patterns: overestimation  
occurs predominantly under high UV conditions, whereas underestimation is more frequent at low to  
moderate radiation levels. The broader scatter at extreme values also suggests that the model's  
performance might be sensitive to solar zenith angles, particularly during sunrise and sunset.

In addition, boxplots of R, MBE and RMSE for all stations combines are shown in Figure 4b. In  
235 particular, the RMSE for the majority of stations ranges between  $3.9 \text{ W m}^{-2}$  and  $5.6 \text{ W m}^{-2}$ , with a mean  
value of  $4.98 \text{ W m}^{-2}$  and a maximum value not exceeding  $9 \text{ W m}^{-2}$ . The mean and median MBE values  
are  $-0.13 \text{ W m}^{-2}$  and  $0.27 \text{ W m}^{-2}$ , respectively, indicating a small overall error. The R-value at most sites  
is greater than 0.9. The median value is notably elevated, approaching the upper quartile, which  
collectively suggests that the overall R-value remains satisfactory and that the model's performance has  
240 strong spatial consistency and robustness, even in regions with varying climatic conditions. Figures 4c  
and 4d display the RMSE and MBE results, respectively, for each station. About half of the stations are

affected by only low RMSEs below  $5 \text{ W m}^{-2}$ . The two most challenging stations are Taihu (TAL) and Aksu (AKA). The Taihu (TAL) station, located in the humid, cloud-prone region of eastern China, exhibits significant errors, with RMSE and MBE reach  $8.82 \text{ W m}^{-2}$  and  $-7.59 \text{ W m}^{-2}$ , respectively. These errors are likely influenced by high humidity and the reflective surface of the Taihu Lake, which complicates the estimation of clear-sky UV radiation. The Aksu (AKA) station, situated in the arid desert region of northwestern China, also stands out with  $\text{RMSE} = 6.40 \text{ W m}^{-2}$ ,  $\text{MBE} = -3.09 \text{ W m}^{-2}$ , and  $R = 0.877$ . The larger errors at AKA may be attributed to several factors, including the high aerosol loading in the region, where mineral dust particles from the surrounding deserts can significantly scatter and absorb UV radiation. Additionally, the reflective properties of the desert surface, extreme climatic conditions, and local meteorological variations could further complicate the accurate estimation of clear-sky UV radiation, leading to the observed discrepancies. These unique atmospheric conditions, combined with the station's extreme climatic environment, likely reduce the model's estimation accuracy. (Wu et al., 2022b).





**Figure 4. Validation results of the estimated hourly clear-sky UV dataset against observations at 37 CERN stations from 2005-2013. (a) Density scatter plot between the estimated dataset and observations. (b) Box plots of the three statistical error metrics (RMSE, MBE, and R). (c) Spatial distribution of RMSE at individual stations. (d) Spatial distribution of MBE at individual stations.**

Figure 5 presents similar results as in Figure 4, but on a daily mean basis. From Figure 5a, the overall validation results are satisfactory, with an  $R$  of 0.907, an  $RMSE$  of  $3.37 W m^{-2}$ , and an  $MBE$  of  $0.14 W m^{-2}$ . The box plots in Figure 5b show that most stations have  $RMSE$  values below  $4 W m^{-2}$ , reflecting consistent performance across regions. The  $MBE$  values are tightly clustered around zero, indicating minimal biases overall. Additionally,  $R$ -values for most stations range between 0.88 and 0.95, demonstrating the model's strong spatial consistency and reliable performance. The spatial distributions of  $RMSE$  and  $MBE$  in Figure 5c and 5d reveal patterns similar to those observed with the hourly estimates. Whereas most stations maintain low  $RMSE$  and near-zero  $MBE$ , the same two stations as before, TAL and AKA exhibit notable discrepancies again. For instance, the daily-mean  $RMSE$ ,  $MBE$ , and  $R$  at AKA are  $7.07 W m^{-2}$ ,  $-6.16 W m^{-2}$ , and 0.783, respectively.



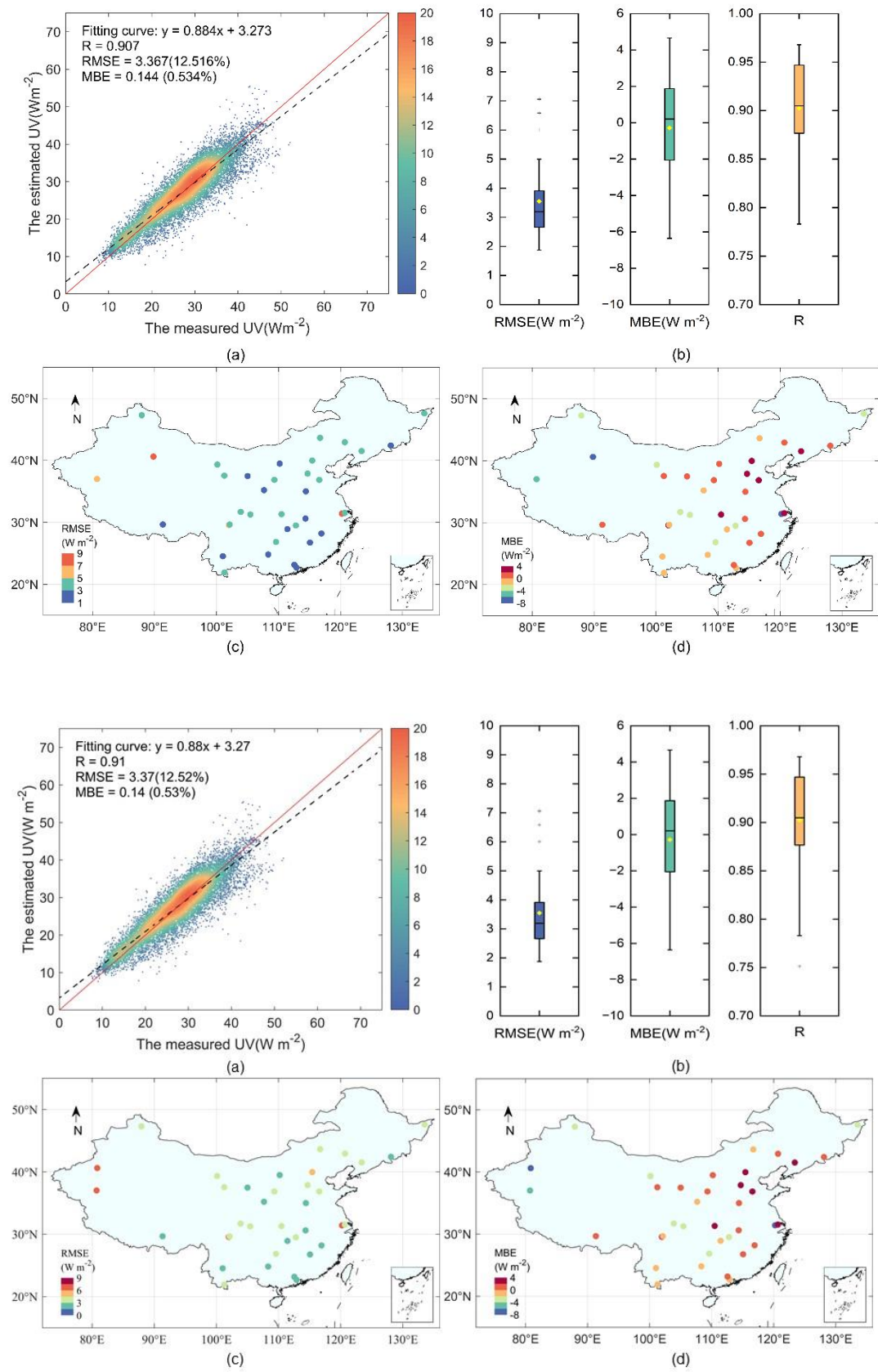
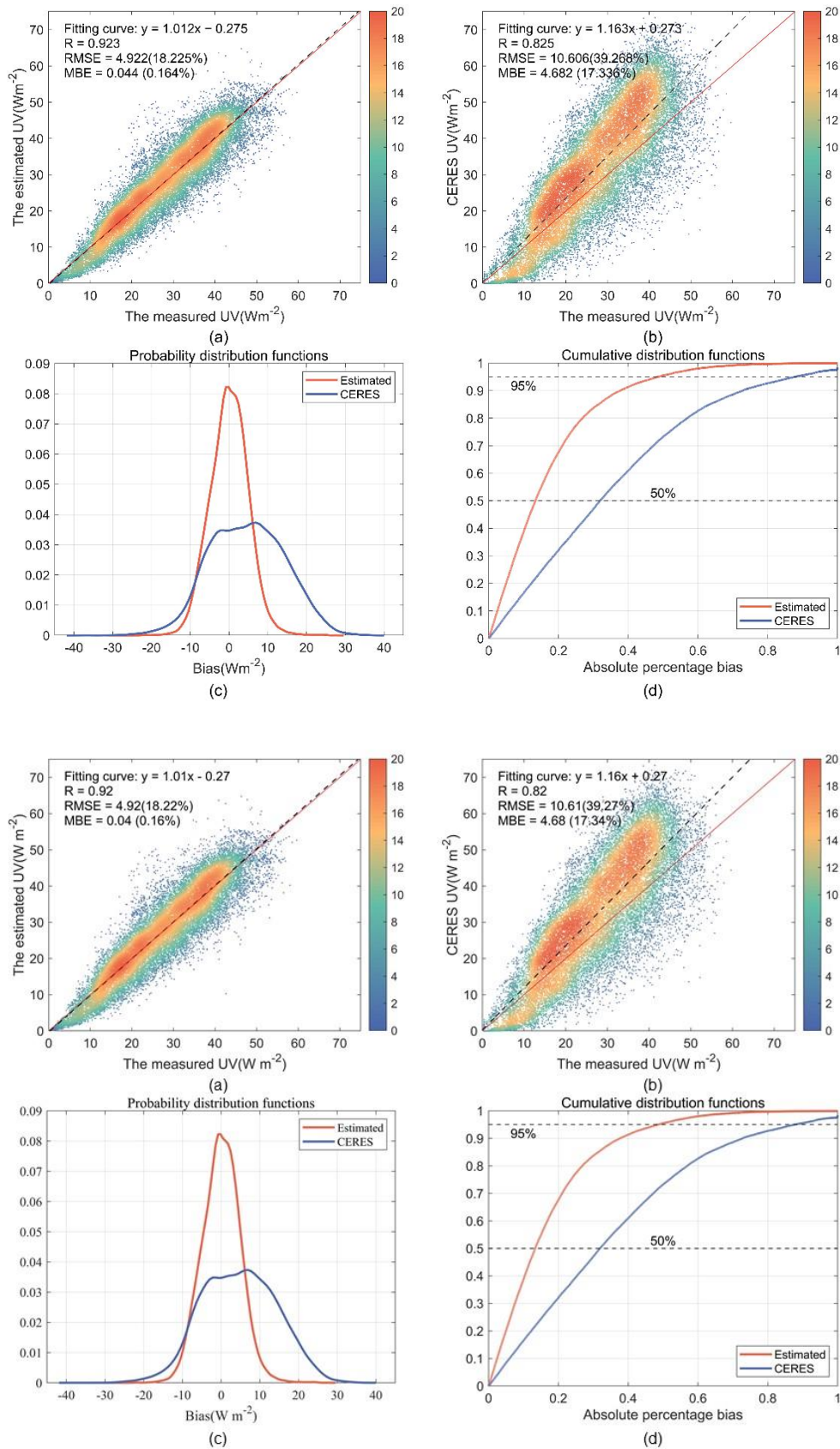


Figure 5. Same as Figure 4 but for daily mean clear-sky UV radiation.

Overall, the validation results demonstrate that the model reliably captures hourly and daily mean  
275 clear-sky UV radiation across diverse environments, maintaining high accuracy and minimal bias at all  
37 sites, with the possible exception of two challenging stations. The ~~normally~~ strong agreement between  
the model estimates and observations underscores its robustness. These findings support the reliability  
of the SMARTS model in estimating clear-sky UV radiation. A supplementary quantile binning analysis  
was conducted to guard against regression bias from uneven data density in our large sample. The results,  
280 shown in Appendix Figure A2, align with Figures 4 and 5, confirming the model's excellent overall  
performance and offering a more robust characterization of its behavior at the extremes.

#### 4.2 Benchmarking the estimated clear-sky UV radiation against the CERES product

A comparison between the present hourly UV estimates and those from CERES SYN1deg is  
desirable as a form of benchmarking, owing to the important status of the latter as a reference global  
285 database. Using CERN observations in 2013, the results in Figure 6 demonstrate that the present UV  
product significantly outperforms SYN1deg in terms of overall accuracy. As shown in Figures 6a and 6b,  
the estimated hourly clear-sky UV radiation achieves an R of 0.923, an RMSE of  $4.92 \text{ W m}^{-2}$ , and an  
MBE of  $0.04 \text{ W m}^{-2}$ , which are substantially better than the CERES product ( $R = 0.825$ ,  $\text{RMSE} = 10.61 \text{ W m}^{-2}$ ,  
 $\text{MBE} = 4.68 \text{ W m}^{-2}$ ). Notably, SYN1deg exhibits systematic biases, with overestimation at high UV  
290 values and underestimation at low UV values, whereas the newly proposed product aligns more closely  
with the observed values. The distribution of the PDF (Figure 6c) further highlights the differences  
between the two products. The bias distribution of the newly proposed product closely resembles a  
Gaussian curve, with a sharp peak near zero, indicating small and stable deviations across most data  
points. In contrast, SYN1deg displays a broader and flatter PDF curve, suggesting larger and more widely  
295 distributed errors, hence lower precision than the new SMARTS-derived product. Figure 6d reinforces  
this conclusion: the CDF curve for ARE is steeper for the newly proposed product, reaching the 95%  
threshold for a much lower mean error than SYN1deg. This confirms that a larger fraction of the  
estimated product's deviations is confined to smaller bias ranges.

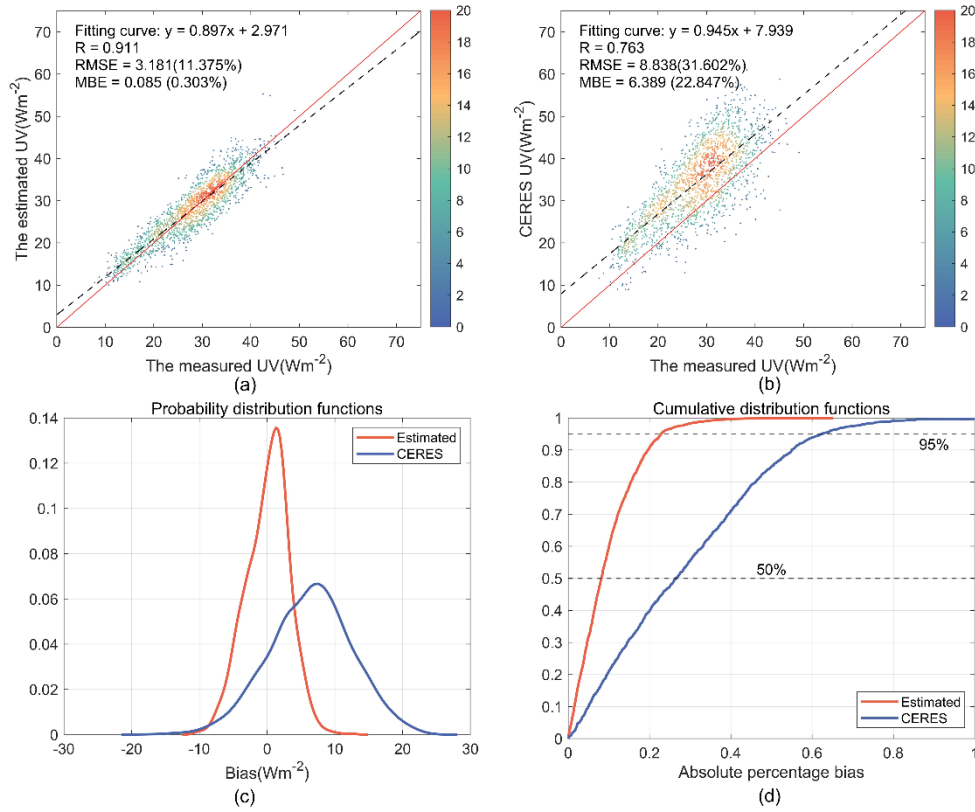


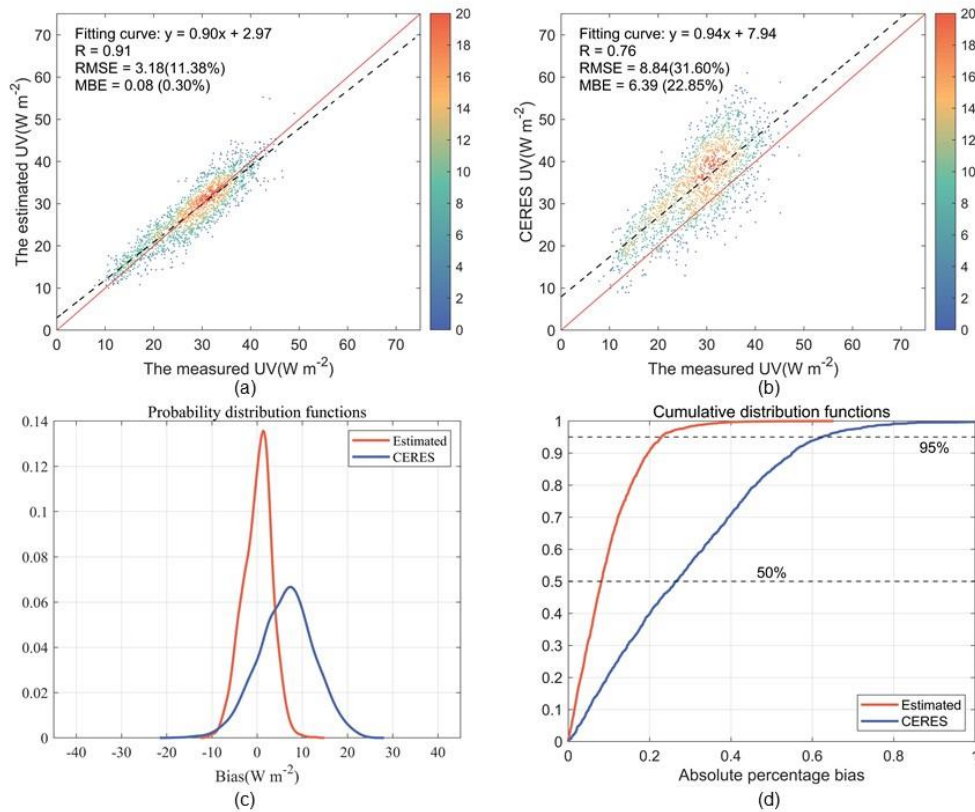
**Figure 6. Hourly assessment results for two clear-sky UV radiation products. (a) Density scatterplot between the hourly clear-sky SMARTS-derived UV irradiance estimates and CERN observations. (b) Same as (a) but**



for the CERES SYN1deg estimates. (c) Probability distribution function of the bias for the SMARTS-derived product (orange line) and the CERES product (blue line); (d) Cumulative distribution function of the absolute relative error for the SMARTS-derived product (orange line) and the CERES product (blue line).

Figure 7 is similar to Figure 6, but on a coarser daily-mean resolution. The R value for the new UV radiation product, 0.911, is markedly higher than the CERES product (0.763). In parallel, the RMSE and MBE for the new product, 3.18 W m<sup>-2</sup> and 0.09 W m<sup>-2</sup>, respectively, are much lower than their CERES counterparts (8.84 W m<sup>-2</sup> and 6.39 W m<sup>-2</sup>). From Figure 7b, it is evident that the CERES product systematically overestimates UV radiation, as indicated by the positive deviation of its best-fit line from the 1:1 line in almost all cases. In contrast, the estimated product aligns much more closely with the observations. The PDF and CDF in Figures 7c and 7d further emphasize the differences between the two products, and confirm the results in Figures 6c and 6d.





**Figure 7. Same as Figure 6 but for daily mean results.**

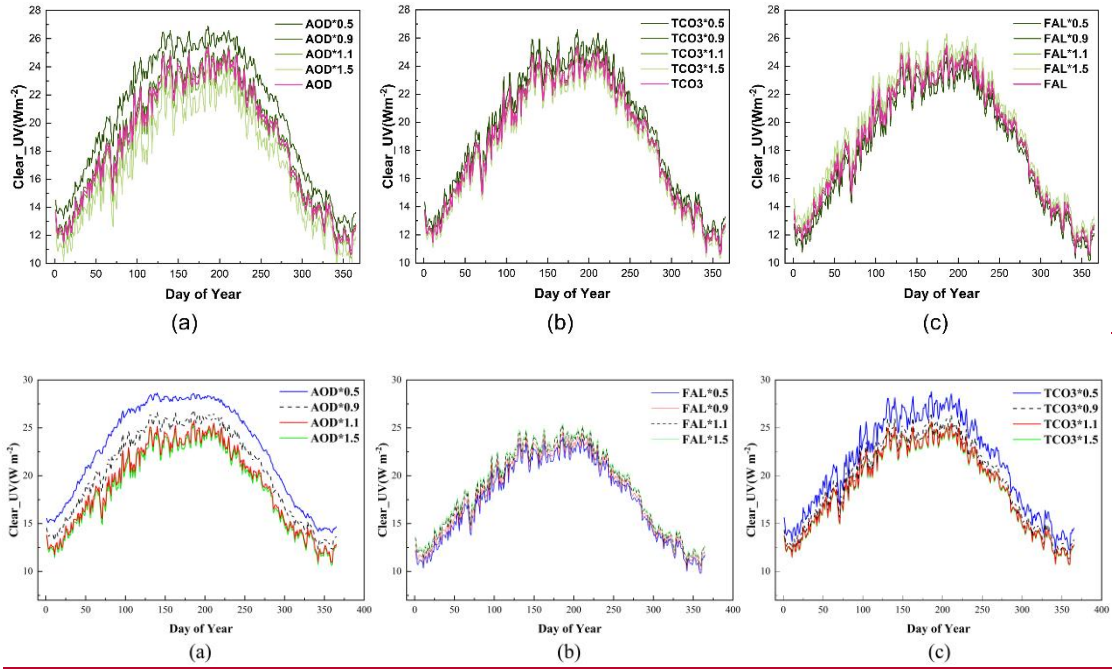
In summary, the hourly and daily mean clear-sky UV radiation product proposed here demonstrates significantly better performance than the CERES SYN1deg product in terms of accuracy, stability, and error distribution. As shown in the binned fitting results in Appendix Figure A3, both products exhibit distribution characteristics similar to those described previously. The CERES SYN1deg product exhibits systematic biases in its representation of UV radiation, overestimating at high values and underestimating at low values. This results in a border and flatter probability density function compared to the new CHUV product, which indicates larger errors and lower overall precision. Consequently, the SYN1deg product also demonstrates a significantly weaker agreement with daily mean observations than the CHUV product.

#### 4.3 Sensitivity of the SMARTS UV predictions to key input variables

The results above indicate that the SMARTS-based UV estimates are satisfactory in general, but are nevertheless affected by reduced accuracy at a few sites where challenging situations are likely to impact the model's inputs. It is thus desirable to evaluate the sensitivity of the modeled UV estimates to discrepancies in the atmospheric inputs. To that effect, a series of scaling factors (0.5, 0.9, 1.1, and 1.5)

are applied to three key input variables as way to analyze the intrinsic model's sensitivity to them. These inputs are the aerosol optical depth (AOD), the total column ozone (TCO3), and the forecast albedo (FAL). In all cases, the clear-sky UV irradiances estimated using the original input data are used as the baseline reference. As shown in Figure 8, the 37-station mean clear-sky UV irradiance thus obtained exhibits a typical unimodal distribution, peaking in summer and reaching a minimum in winter, as could be expected. This seasonal variation is closely tied to changes in solar zenith angle. Among the four input parameters tested here, AOD induces the most significant impact on clear-sky UV radiation (Figure 8a). As AOD increases, the UV irradiance decreases substantially, highlighting the strong attenuation effect of aerosols in the UV ~~(from -50.06% to -40.90% per AOD unit)~~. The sensitivity analysis revealed a strong negative relationship between AOD and surface UV irradiance. Specifically, a one-unit increase in AOD was associated with a fractional reduction in UV irradiance ranging from -50.06% to -40.90%. This is consistent with the physical properties of aerosols, which scatter and absorb radiation mostly in the UV spectrum. As TCO3 increases incrementally relative to its original value, the clear-sky UV radiation decreases progressively (Figure 8b), which reflects the strong absorption of ozone in the UV spectrum, although this is confined to wavelengths below  $\approx 340$  nm only. Surface albedo (FAL) is shown to enhance UV radiation when it increases (Figure 8c). Higher albedo values, such as those from snow-covered surfaces, amplify the reflection of UV radiation back into the atmosphere, which can subsequently increase the downward UV irradiance through the atmospheric backscattering effect (Shine et al., 2012).

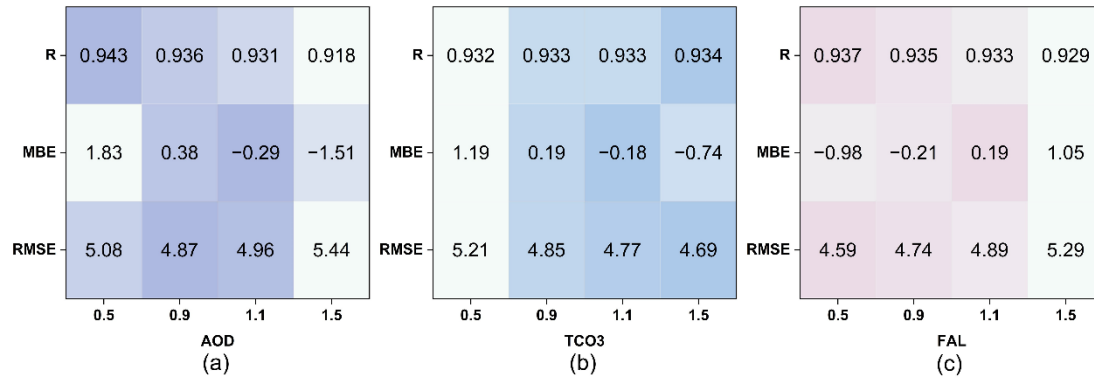
A key limitation in this study arises from the selection of the aerosol model in the SMARTS radiative transfer simulations. To maintain consistency with the continental focus of our work, the built-in "Continental" aerosol model was employed. While this model is representative of typical inland aerosol species such as sulfates, nitrates, black carbon, and dust, it does not include an explicit parameterization for sea salt aerosols. This discrepancy could introduce systematic biases in simulated UV radiation in coastal regions and inland areas subject to marine air mass advection, as sea salt aerosols differ significantly from continental types in their size distribution, and scattering efficiency (Zhu et al., 2022; Kouvarakis et al., 2002; Chatzopoulou et al., 2025). Future research should aim to integrate a more sophisticated hybrid or maritime aerosol model and couple it with higher-resolution aerosol reanalysis or observational data to better constrain the aerosol-type dependence and improve the accuracy of UV estimates in complex coastal-continental transition zones.



**Figure 8. Mean annual variation of estimated clear-sky UV irradiance at 37 CERN stations under different conditions. The baseline (red curve) uses the original input data. The four green curves represent the estimates obtained after scaling the indicated input.**

Figure 9 presents the validation results of the SMARTS model's clear-sky UV radiation estimates against station observations under different scaling factors. For AOD, as the scaling factor increases from 0.5 to 1.5, the R value decreases from 0.943 to 0.918, indicating a slight decline in the model's correlation with observations. The MBE shifts from  $1.83 \text{ W m}^{-2}$  to  $-1.51 \text{ W m}^{-2}$ , suggesting that the model overestimates UV radiation at lower scaling factors and underestimates it at higher scaling factors. The RMSE initially decreases from  $5.08 \text{ W m}^{-2}$  to  $4.87 \text{ W m}^{-2}$  (scaling factors 0.5–0.9), and subsequently increases to  $5.44 \text{ W m}^{-2}$  when the scaling factor exceeds 1, demonstrating a reduction in the model's predictive capability with increasing AOD scaling factors. As the scaling factor for TCO3 increases, the R value remains stable between 0.932 and 0.934, showing minimal variation. The MBE decreases from  $1.19 \text{ W m}^{-2}$  to  $-0.74 \text{ W m}^{-2}$ , indicating a transition from overestimation to underestimation. Meanwhile, the RMSE decreases from  $5.21 \text{ W m}^{-2}$  to  $4.69 \text{ W m}^{-2}$ , suggesting that the model's prediction error is reduced at higher TCO3 scaling factors. When the FAL scaling factor reaches 1.5, the model exhibits the lowest overall accuracy, with an R value of 0.929, MBE of  $1.05 \text{ W m}^{-2}$ , and RMSE of  $5.29 \text{ W m}^{-2}$ , indicating that increasing FAL scaling factors significantly amplify the model's prediction error. The MBE shifts from  $-0.98 \text{ W m}^{-2}$  to  $1.05 \text{ W m}^{-2}$ , demonstrating that the model underestimates at lower scaling factors and overestimates at higher scaling factors. These results reveal the complex influence of

different input parameters on UV radiation estimation and provide critical insights for optimizing model performance.



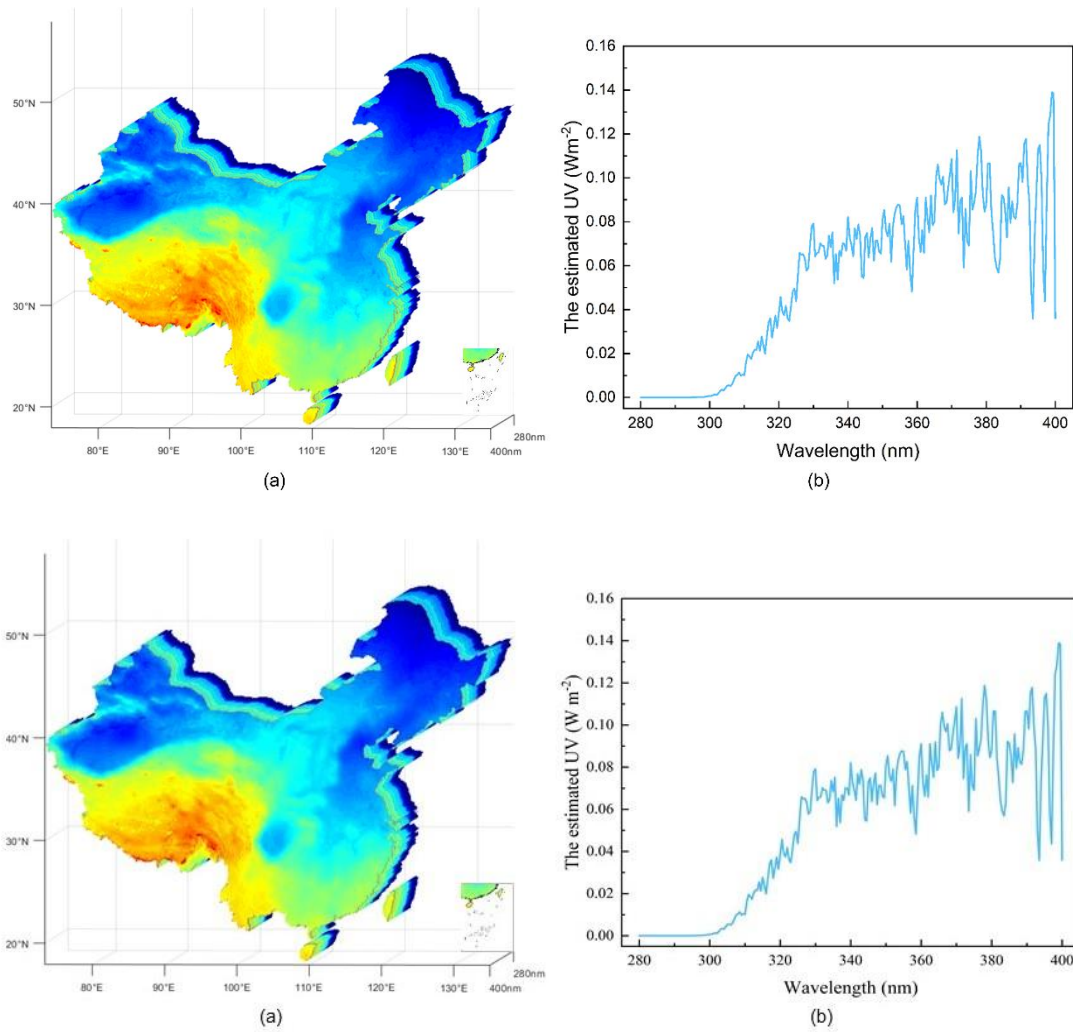
**Figure 9. Accuracy validation of clear-sky UV radiation estimates and station observations under different conditions.**

#### 4.4 Evaluating the spatial and temporal variability of clear-sky UV radiation

##### 4.4.1 Evaluating the spatial variability of clear-sky UV radiation across wavelengths

In this study, we generated a stereogram of clear-sky UV radiation by wavelength using the SMARTS model. The data range spans 280–400 nm with a step size of 0.5 nm, encompassing radiation values across 241 wavelength bands (Figure 10a). The first layer of this figure illustrates the distribution of total UV radiation under clear-sky conditions across China in 2023, while the subsequent 241 layers represent the radiation value distributions for individual wavelength bands. Figure 10b shows an unfolded representation of the mean clear-sky UV radiation values across China for each wavelength bands from Figure 10a, showing a gradual increase in UV radiation intensity with increasing wavelength.

In the short wavelength range (280–310 nm), the radiation is relatively weak, likely due to strong absorption by the ozone layer. In contrast, the long wavelength range (near 380–400 nm) exhibits stronger radiation, indicating greater penetration capability. By analyzing the radiation intensity at different wavelengths, the impact of UV radiation at various wavelengths on human health and the environment can be studied.

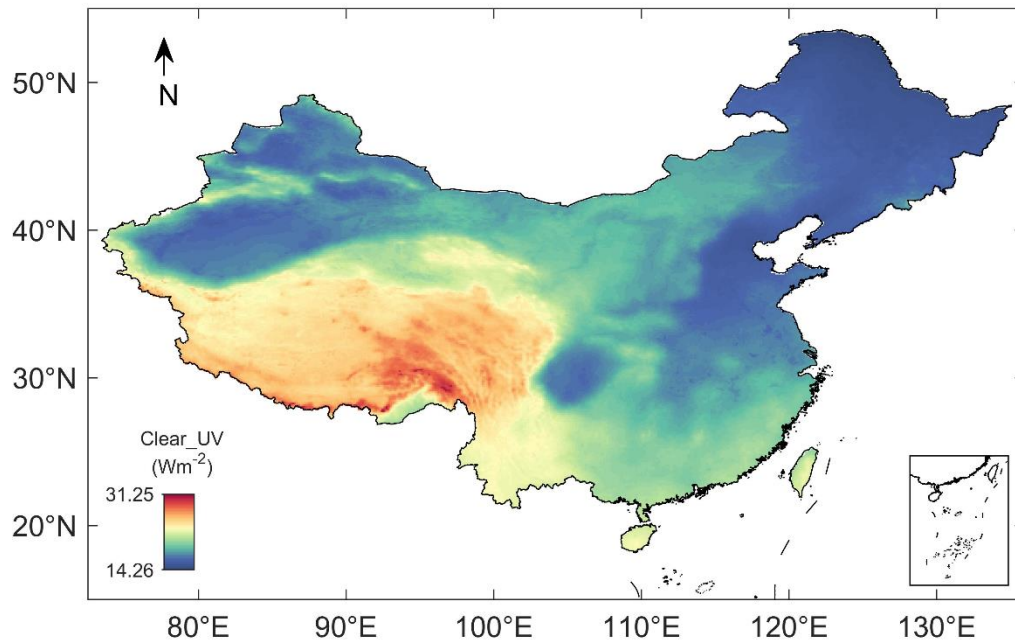


**Figure 10. (a) 3D hyperspectral distribution of clear-sky UV radiation in China (2023). (b) Spectral variation of clear-sky UV radiation (280–400 nm) across entire China.**

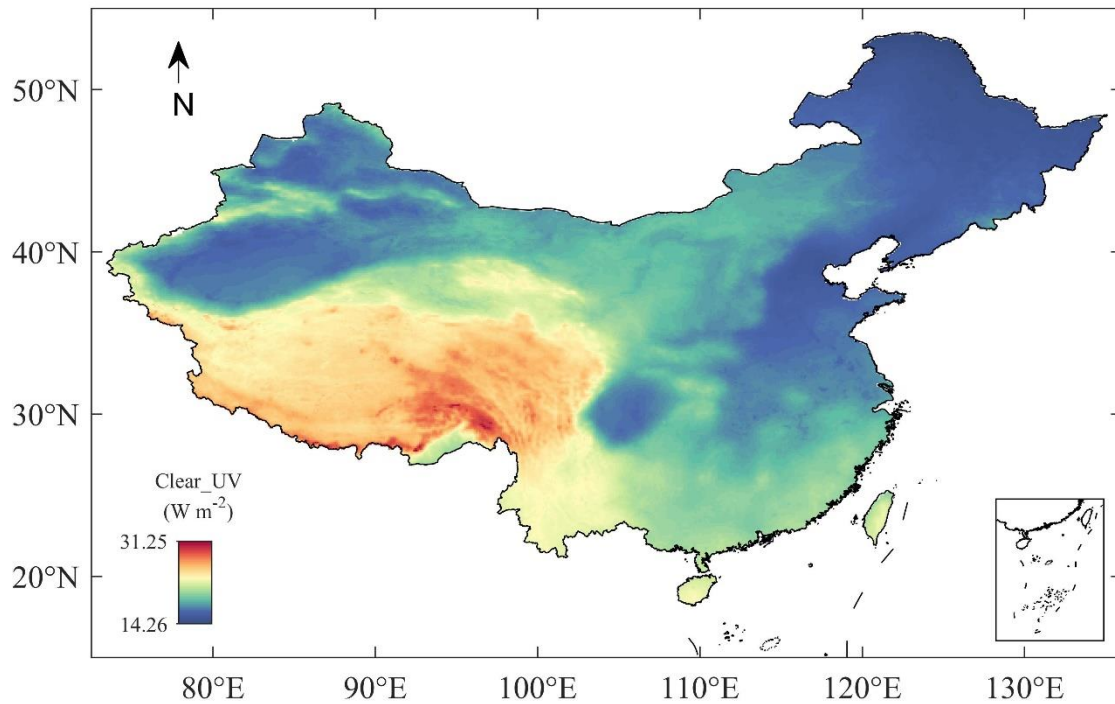
#### 4.4.2 Evaluating the spatial variability of clear-sky UV radiation

The SMARTS model was employed to reconstruct a long-term dataset of clear-sky UV radiation across China from 1981 to 2023. Figure 11 presents the annual average spatial distribution of clear-sky UV radiation in China during 1981-2023. The clear-sky UV radiation values range from  $14.26 W m^{-2}$  to  $31.25 W m^{-2}$ , with an overall annual average value of  $20.05 W m^{-2}$ . Spatially, clear-sky UV radiation exhibits a distinct west-to-east decreasing gradient, accompanied by lower values in northeastern China compared to the southern regions. The Tibetan Plateau, characterized by its high altitude, exhibits the highest clear-sky UV radiation levels in the country. The reduced atmospheric thickness over high-altitude regions results in less attenuation of UV radiation, thereby enhancing surface UV levels. In contrast, northeastern China experiences lower UV radiation due to its higher latitude, where lower solar

altitude angles lead to greater atmospheric absorption and scattering, significantly attenuating surface  
415 UV radiation. Additionally, the Sichuan Basin is identified as a region with notably low clear-sky UV  
radiation levels. This is attributed to its low elevation and surrounding mountains, which limits the  
amount of solar radiation received, which enhances UV radiation absorption and scattering within the  
atmosphere (Qin et al., 2020a).



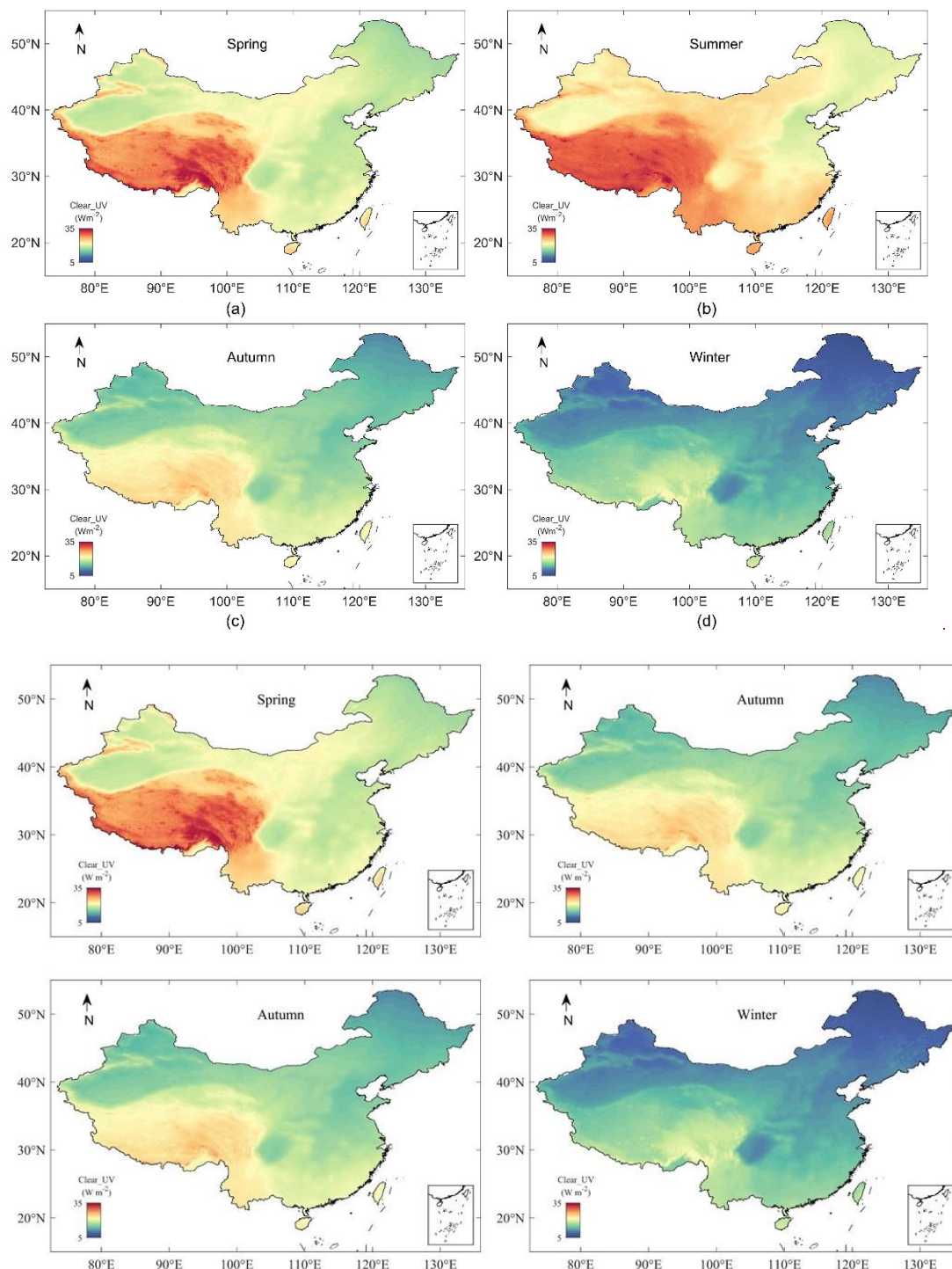




**Figure 11. Annual average spatial distribution of clear-sky UV radiation in China from 1981 to 2023.**

As illustrated in Figure 12, the clear-sky UV radiation in China exhibits pronounced seasonal variations, primarily driven by changes in the solar zenith angle and maximum sunshine duration throughout the year. The highest UV radiation levels are observed in summer (Figure 12b), particularly over the Tibetan Plateau, where values approach  $35 \text{ W m}^{-2}$ . Conversely, the winter (Figure 12d) demonstrate the least intensity of UV radiation, with values in northern regions reaching only  $5\text{-}10 \text{ W m}^{-2}$ . Spring (Figure 12a) and autumn (Figure 12c) exhibit intermediate radiation levels, reflecting the seasonal transition between the extremes of summer and winter. The average national UV radiation levels experienced a notable increase of  $0.8962 \text{ W m}^{-2}$  in the spring months when compared to the annual mean. Conversely, during the autumn season, a decline of approximately  $0.9277 \text{ W m}^{-2}$  was observed. From a spatial perspective, the Tibetan Plateau consistently records the highest UV radiation levels across all seasons due to its ~~considerable~~ high-altitude (mean elevation exceeding 4000 meters) and thinner atmosphere. In contrast, eastern regions of China, characterized by lower altitudes and denser atmospheric layers, display comparatively lower UV radiation levels. These spatial patterns underscore the influence of both altitude and latitude on the distribution of UV radiation across China.





**Figure 12. Average seasonal spatial distribution of clear-sky UV radiation in China from 1981 to 2023.**

#### 4.4.3 Evaluating the inter-annual variability and trends of clear-sky UV radiation

440 The annual trend of clear-sky UV radiation in the Chinese region from 1981 to 2023 is shown in Figure 13. Over the 43-year period, the annual average clear-sky UV radiation demonstrates a slight overall upward trend ( $+0.0237 \text{ W m}^{-2} \text{ yr}^{-1}$ ), with annual mean values ranging from  $16.02 \text{ W m}^{-2}$  to  $21.33$

W m<sup>-2</sup>. The analysis of trends in long-term environmental time series is recognized as being sensitive to the selected data range and to methodology of detrending (Wu et al., 2007). In this study, t

445 annual variability can be divided into three distinct phases. From 1981 to 1994, clear-sky UV radiation exhibited a decreasing trend ( $-0.0488 \text{ W m}^{-2} \text{ yr}^{-1}$ ). This reduction is primarily attributed to the eruption of the El Chichón volcano in Mexico in 1982 and the Mount Pinatubo eruption in the Philippines in 1991.

A large amount of smoke and volcanic ash was released from the stratosphere into the troposphere, leading to a sharp increase in aerosol particle concentration in the atmosphere. High concentrations of

450 aerosols significantly reduce the intensity of UV radiation reaching the Earth's surface through scattering and absorption processes. (Fang et al., 2021). In 1992, clear-sky UV radiation reached its lowest value in 43 years at  $16.02 \text{ W m}^{-2}$ . After 1992, the impact of volcanic eruptions on clear-sky UV radiation gradually diminished, with the overall trend levelling off by 1995. Following this period, from 1995 to

2009, with rapid economic development and the extensive burning of fossil fuels, air pollution became

455 more severe, leading to an increase in anthropogenic aerosols. At the same time, the rise in ozone concentration over China enhanced the scattering and absorption processes in the atmosphere (Verstraeten et al., 2015), thereby reducing the UV radiation reaching the Earth's surface ( $-0.0329 \text{ W m}^{-2} \text{ yr}^{-1}$ ).

Since 2010, clear-sky UV radiation has risen sharply ( $+0.0670 \text{ W m}^{-2} \text{ yr}^{-1}$ ), ~~coinciding~~

~~with~~following the implementation of stringent air pollution control policies in China ~~that, which~~

460 ~~effectively led to a reductioning the concentration of in atmospheric~~ aerosols, ozone, and other pollutants

~~in the atmosphere by through~~ curbing coal consumption and industrial emissions (He et al., 2016). The

continuous improvement in air quality has significantly enhanced the intensity of UV radiation under

clear-sky conditions, showing a significant upward trend. Finally, while the identified trend is statistically

significant and physically consistent, it should be noted that its quantitative value is methodology- and

465 period-dependent, a common consideration in non-stationary time series analysis.

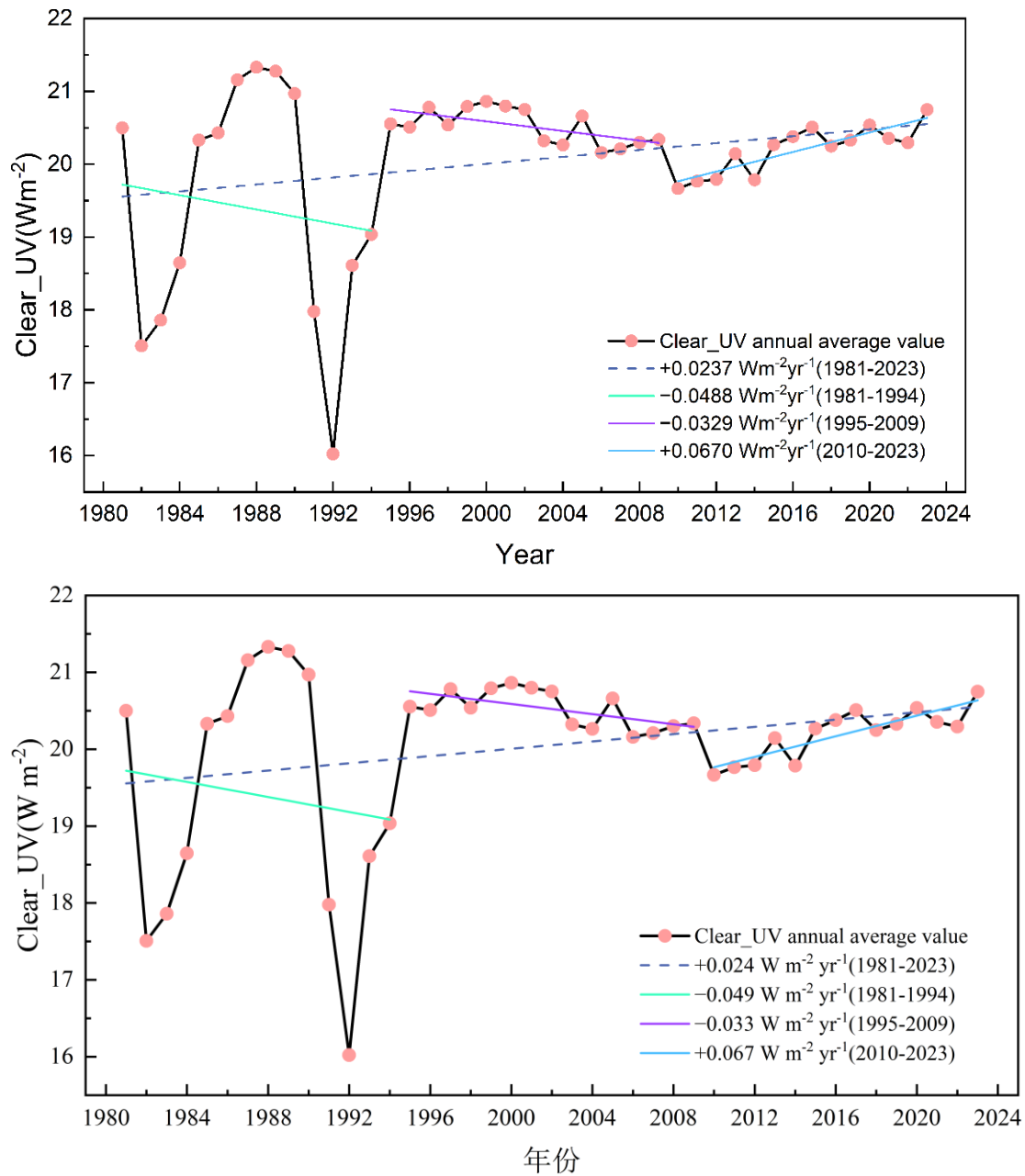


Figure 13. Inter-annual variation trend of clear-sky UV radiation in China, 1981-2023.

## 5 Data availability

470 The hourly clear-sky UV radiation dataset reconstructed in this study, covering 43 years (1981–  
 2023), has been uploaded to figshare, and is stored in NetCDF format. Users can access the complete  
 dataset via the links: <https://doi.org/10.6084/m9.figshare.28234298> (Qi et al., 2025). The one-year  
 datasets include 365 or 366 NetCDF files, with file names following the format “UV\_yyyymmdd”, where  
 “yyyy” represents the year, “mm” represents the month, and “dd” stands for the day. The unit of the data  
 475 is  $\text{W m}^{-2}$ . Each file contains three variables: solar UV radiation, longitude, and latitude, with dimensions

of  $616 \times 356 \times 24$ . The dataset covers the region from  $73.5^\circ\text{E}$  to  $135^\circ\text{E}$  and  $18^\circ\text{N}$  to  $53.5^\circ\text{N}$ , with a spatial resolution of  $0.1^\circ$  (approximately 10 km). The data is presented in local standard time (LST), corresponding to UTC+8 for China.

In addition, leveraging the spectral computation capabilities of the SMARTS model, the dataset has  
480 been further refined to provide UV radiation data at 0.5 nm intervals within the wavelength range of 280–400 nm, covering a total of 241 bands. This refined dataset has also been uploaded to figshare (<https://doi.org/10.6084/m9.figshare.28234298>) (Qi et al., 2025), accessible through their respective links for download. Each file is named using the format “UV\_05nm\_yyyymmddhh,” where “yyyy” denotes the year, “mm” represents the month, “dd” stands for the day, and “hh” denotes the hour, with  
485 units in  $\text{W m}^{-2}$ . Each file contains four variables (solar UV radiation, longitude, latitude, and wavelength), with dimensions of  $616 \times 356 \times 241$ .

## 6 Summary and conclusions

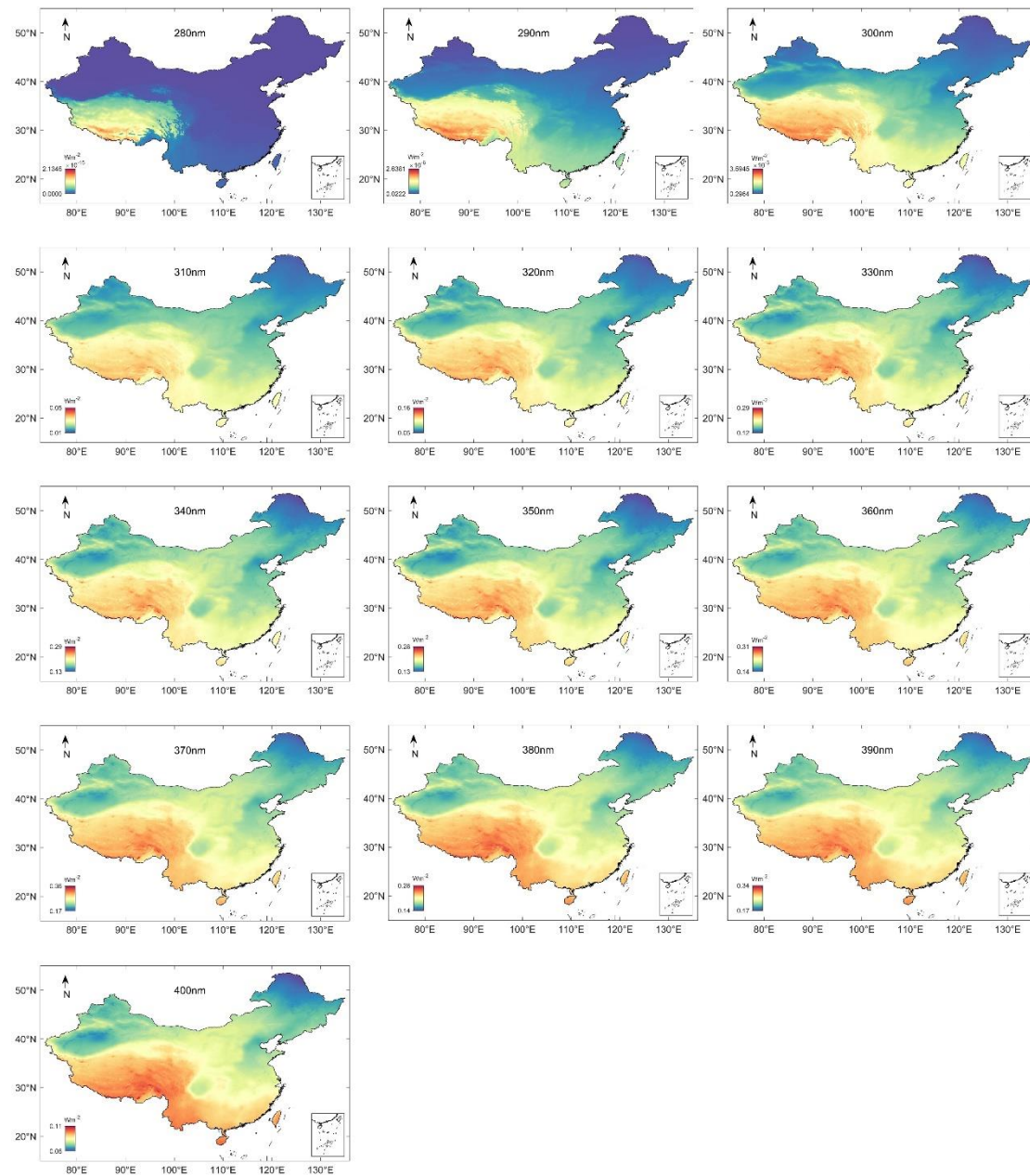
We used the SMARTS model to reconstruct a long-term (1981-2023) hourly clear-sky UV radiation dataset ( $10 \text{ km} \times 10 \text{ km}$ ) in China. The SMARTS model is capable of generating wavelength-specific  
490 clear-sky UV radiation with a spectral resolution of 0.5 nm, covering the 280-400 nm range, which enhances its ability to analyze spectral UV characteristics. Key inputs to the model include conventional meteorological products from ERA5 (e.g., relative humidity, total column ozone, total column water vapor, surface pressure, 2m temperature and forecast albedo) and aerosol properties from MERRA2. The generated clear-sky UV radiation dataset was rigorously validated against CERN site observations and  
495 further compared with UV radiation estimates from CERES products. For the hourly clear-sky UV radiation, the overall R, RMSE and MBE were 0.919,  $5.07 \text{ W m}^{-2}$  and  $-0.07 \text{ W m}^{-2}$ , respectively. On the daily scale, these values were 0.907,  $3.37 \text{ W m}^{-2}$  and  $0.14 \text{ W m}^{-2}$ , respectively. The generated clear-sky UV radiation dataset significantly outperforms the CERES product on both temporal scales.

In addition, this study conducted a sensitivity analysis on the key input parameters (AOD, TCO3,  
500 and FAL) of the SMARTS model using scaling factors of 0.5, 0.9, 1.1, and 1.5. The results indicate that AOD and TCO3 are the primary attenuating factors for UV radiation, while FAL has an enhancing factor. Specifically, as the AOD scaling factor increases, the model's predictive capability significantly declines. At higher TCO3 scaling factors, the model's prediction error decreases. Conversely, increasing the FAL

scaling factor substantially amplifies the model's prediction error. The spatial distribution of clear-sky  
505 UV radiation in China generally exhibits a pattern of "higher in the west and lower in the east, higher in  
the south and lower in the north". In addition, there is significant seasonal variation. The annual mean  
values of clear-sky UV radiation range from  $14.26 \text{ W m}^{-2}$  to  $31.25 \text{ W m}^{-2}$ , with the overall annual mean  
reaching  $20.05 \text{ W m}^{-2}$ . By analyzing the spectral UV radiation distribution, it was observed that the clear-  
sky UV radiation increases with wavelength across the 280-400 nm range. Throughout the study period,  
510 clear-sky UV radiation exhibited a slight increasing trend ( $+0.0237 \text{ W m}^{-2} \text{ yr}^{-1}$ ).

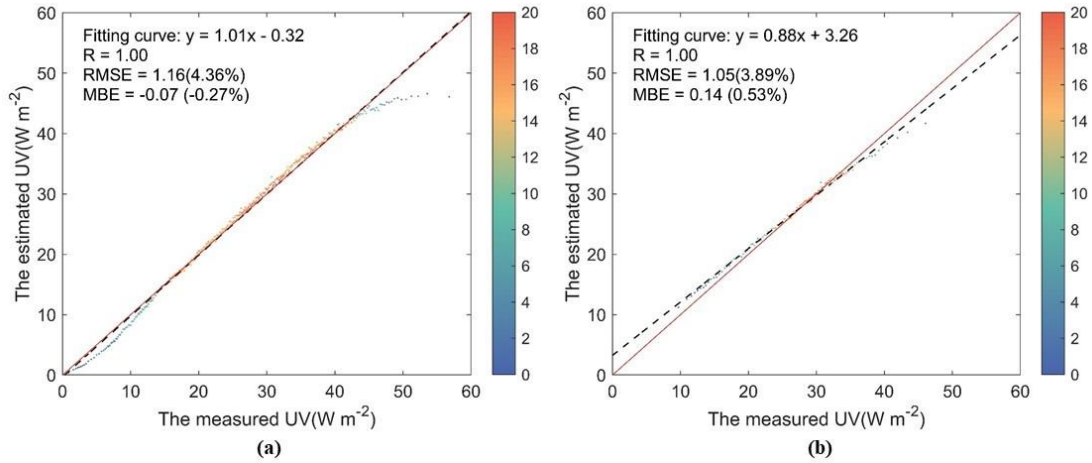
This study reconstructed a long-term, high-temporal, and high-spatial resolution dataset of clear-  
sky UV radiation and further refined it into a spectral UV radiation dataset with 0.5 nm intervals. We  
anticipate that this dataset will provide valuable support for research in fields such as human health and  
ecological environments. However, as the SMARTS model only simulates UV radiation under clear-sky  
515 conditions, future research should integrate the effects of cloud cover to develop an all-sky UV radiation  
dataset. Furthermore, the scope of the study could be expanded from the national to the global scale,  
enabling a more comprehensive spatial and temporal analysis.

## Appendix A: Additional figures

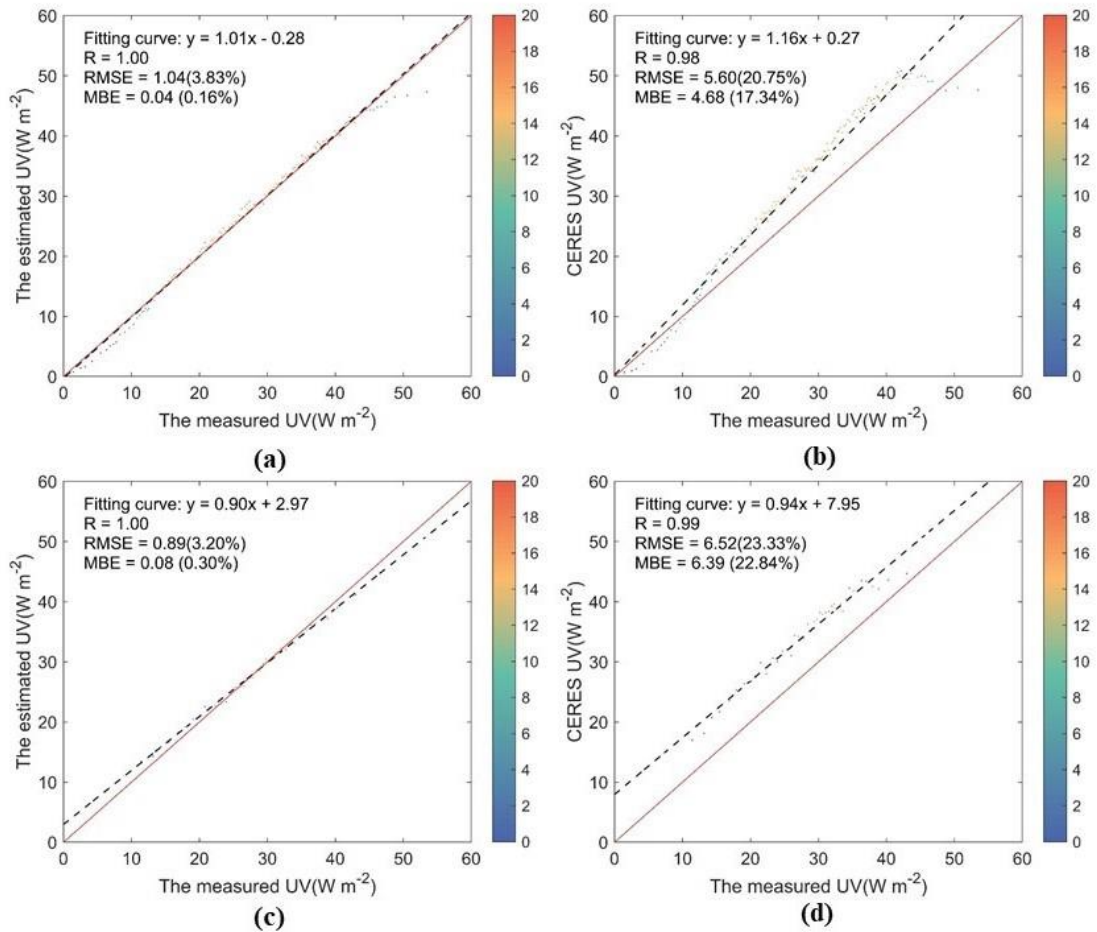


520 Figure A1. Spatial distribution map of specific spectral bands in China (2023)





**Figure A2. Binning scatter validation results of the estimated clear-sky UV dataset against observations at 37 CERN stations from 2005-2013. (a) Density scatter plot between the hourly estimated dataset and observations; (b) same as (a) but for daily results.**



**Figure A3. Binning scatterplot of assessment results for two clear-sky UV radiation products. (a) Density scatterplot between the hourly clear-sky SMARTS-derived UV irradiance estimates and CERN observations. (b) Same as (a) but for the CERES SYN1deg estimates. (d) and (e) same as (a) (b) but for daily results.**

### **Author contributions**

530 QQ: designed the research, analyzed the results, wrote and revised the manuscript. YT: calculated the dataset and validation. CAG: support the initial code of SMARTS model. WQ: designed the research, revised the manuscript. HB: revised the manuscript and provide the CERN Data. TS: analyzed the data. MZ: revised the manuscript. MW: revised the manuscript. LW: revised the manuscript.

### **Competing interests**

535 The authors have the following competing interests: Martin Wild is a member of the editorial board of Earth System Science Data.

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545 websites (<https://cds.climate.copernicus.eu> and <https://disc.gsfc.nasa.gov>). The authors would like to thank the staff of the data management and production organizations for their valuable work.

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