



1	Multi-Element Dataset Across Diverse Climatic Zones and Soil Profiles in China's
2	Mountains
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16 Abstract

17	Mountain ecosystems are crucial for global biodiversity conservation and climate regulation, yet their
18	response to environmental change remains poorly understood due to limited high-resolution, multi-
19	element datasets. Here, we present a comprehensive geochemical dataset comprising more than 1,300
20	soil samples collected from 166 sites across 30 mountain regions in China, spanning five major climatic
21	zones and representative vegetation types. Soil samples were systematically collected from three
22	standardized horizons (organic, surface mineral, and parent material), and analyzed for the concentrations
23	of 24 elements, including macronutrients (e.g., phosphorus, potassium, calcium, magnesium),
24	micronutrients (e.g., iron, molybdenum, manganese, copper), and trace metals (e.g., cadmium, chromium,
25	lead, antimony). To support integrated Earth system analyses, the dataset is accompanied by key site-
26	specific environmental variables, including climate parameters (temperature, precipitation, aridity index),
27	normalized difference vegetation index, soil physicochemical properties (pH, moisture, bulk density),
28	atmospheric nitrogen deposition, and chemical weathering index. The dataset reveals significant vertical
29	stratification in element distributions, with organic horizon enriched in biogenic elements, and deeper
30	horizons dominated by lithogenic components. Spatial patterns along latitudinal, longitudinal, and
31	altitudinal gradients underscore the influence of climate and geology on soil chemistry. This open-access
32	dataset provides a valuable resource for parameterizing and validating biogeochemical models, assessing
33	soil quality in mountain regions, and improving predictions of ecosystem responses to global change.
34	The dataset can be accessed via https://doi.org/10.11888/Terre.tpdc.302620 (Wu et al., 2025b).

35 Key words: Mountain ecosystems, Soil dataset, Elementome, Soil profile, Multi-element





36 1 Introduction

37 Mountain ecosystems, recognized globally as hotspots of biodiversity and critical regulators of Earth's 38 biogeochemical cycles, are under increasing pressure from accelerating global changes (Antonelli et al., 39 2018; Wang et al., 2022) and anthropogenic impacts. The complex topography of mountain regions 40 creates steep climatic and edaphic gradients over short spatial distances, generating distinct 41 biogeochemical niches that are highly sensitive to environmental change (Dainese et al., 2024; Nogués-42 Bravo, 2007). Soils in these regions are central to ecosystem functioning, supporting primary productivity, nutrient cycling, and carbon sequestration (Cui et al., 2022; Sundqvist et al., 2013). However, our 43 44 understanding of how these functions are regulated by the distribution and interaction of multiple soil elements remains limited, particularly under conditions of rapid global change (Kaspari and Powers, 45 46 2016; Tian et al., 2019; Wu et al., 2025a). A fundamental constraint hindering progress in this field is the 47 scarcity of comprehensive, large-scale datasets that simultaneously quantify the distribution and 48 interaction of a wide range of soil elements across diverse mountain landscapes. Overcoming this data 49 deficiency is essential for unraveling the emergent properties of coupled biogeochemical cycles in these 50 vulnerable ecosystems and for developing robust predictive models for their future integrity and 51 resilience.

52 While the roles of key elements such as carbon (C) and phosphorus (P) in ecosystem processes are 53 well established (Elser et al., 2010; Zuo et al., 2024), recent research highlights the importance of a 54 broader suite of elements, including macronutrients, micronutrients, and trace metals, in shaping 55 biogeochemical dynamics (Han et al., 2011; Vallicrosa et al., 2022). For instance, calcium (Ca), 56 magnesium (Mg), and potassium (K) are essential for plant physiological processes (Fernández-Martínez 57 et al., 2021; Peñuelas et al., 2019), while micronutrients such as molybdenum (Mo), copper (Cu), and 58 manganese (Mn) function as critical cofactors in enzymatic pathways driving nitrogen fixation, methane 59 oxidation, and organic matter decomposition (Dai et al., 2023; Hay et al., 2023). Furthermore, the 60 intricate interplay between trace element availability and the structure and function of microbial 61 communities is increasingly recognized as a key control on soil biogeochemistry (Giovannelli, 2023; 62 Shafiee et al., 2021; Zhao et al., 2020). Conversely, elevated concentrations of toxic elements, including 63 aluminum (Al), cadmium (Cd), chromium (Cr), and lead (Pb), can disrupt ecosystem integrity and impair 64 ecosystem health through food chain transfer (Bing et al., 2016; Lynch and St. Clair, 2004; Nagajyoti et





al., 2010). Capturing the distribution and interactions of this wide range of elements is therefore
indispensable for a mechanistic understanding of mountain ecosystem functioning (Fernández-Martínez
et al., 2021; Kaspari and Powers, 2016; Vallicrosa, 2022).

68 Emerging evidence suggests that the composition and distribution of soil elements are regulated not 69 only by individual environmental drivers such as climate, vegetation, and parent material, but also by 70 complex interactions among these factors (Augusto et al., 2017; Haghverdi et al., 2019; Molina et al., 71 2024; Moreno-Jiménez et al., 2022). These interactions can lead to non-linear responses and threshold 72 effects that are difficult to predict without comprehensive, spatially explicit datasets (Feng et al., 2024). 73 Vertical stratification of elements along soil profiles further reflects the combined influence of biological 74 inputs, weathering processes, and leaching losses, offering insights into long-term ecosystem 75 development and nutrient cycling (Cronan et al., 2018; Kirkby, 2018; Jobbágy and Jackson, 2001; 76 Steinnes and Lierhagen, 2018). For instance, biologically cycled elements tend to accumulate in organic-77 rich layers, while lithogenic elements are often more concentrated in deeper mineral layers or parent 78 material (Jobbágy and Jackson, 2004; Agnan et al., 2019; Sayer et al., 2020). However, few studies have 79 captured this vertical dimension across broad environmental gradients, and most existing datasets are 80 limited in geographic scope, sampling depth, or the range of elements analyzed (Moreno-Jiménez et al., 81 2019; Ochoa-Hueso et al., 2023; Shangguan et al., 2014).

82 China's extensive and environmentally diverse mountain systems, which cover over two-thirds of 83 the nation's terrestrial area, offer an exceptional natural laboratory for dissecting the complex interactions 84 among climate, vegetation, geology, and soils that shape global biogeochemical cycles (Han et al., 2011). 85 These mountains span a wide range of climatic zones, from tropical and subtropical rainforests to 86 temperate forests, boreal forests, and alpine tundra, nearly including the full spectrum of Earth's 87 terrestrial biomes (Kou et al., 2021; Lu et al., 2018). The combination of steep altitudinal gradients, varied 88 parent materials, and diverse vegetation types results in a mosaic of soil environments that mirrors the 89 global diversity of pedogenic conditions. Critically, many of these mountain regions remain relatively 90 undisturbed by direct anthropogenic activities, providing a unique opportunity to assess natural controls 91 on soil geochemistry and to establish baselines for detecting future environmental change. Here, we 92 conducted a large-scale soil sampling campaign across 30 mountain ecosystems in China. These sites 93 represent five major climatic zones and typical vegetation types in each region. A total of 1,314 soil





94 samples were collected from three standardized soil horizons (organic (O), surface mineral (A), and 95 parent material (C) horizons). Each sample was analyzed for the concentrations of 24 elements, including 96 macronutrients, micronutrients, and trace metals. In addition, the dataset was integrated with a suite of 97 ancillary variables, including satellite-derived vegetation indices, soil physicochemical properties, and 98 climatic parameters. This dataset provides a rare, high-resolution view of multi-element distributions 99 across spatial and vertical gradients in mountain soils. It enables cross-disciplinary analyses that link 100 geochemical, ecological, and climatic processes, and serves as a critical resource for improving 101 biogeochemical modeling, evaluating soil quality and ecosystem health, and understanding the resilience 102 of mountain systems under global change.

103 2 Materials and methods

104 2.1 Study area

105 Soil samples were collected from 30 mountains located within accessible national and provincial nature 106 reserves across China (Fig. 1). The study area spans a wide geographical range (18.9°N-53.5°N, 101.0°E-107 129.6°E, altitude: 210-4225 m above sea level). These sites represent diverse ecosystem zones, including 108 tropical, subtropical, warm temperate, temperate, and cold temperate regions. Vegetation types include 109 broadleaf forests, mixed broadleaf-coniferous forests, coniferous forests, and shrublands (Bing et al., 110 2021; Cui et al., 2022). The mean annual temperature (MAT) in the study area ranges between -5°C and 111 21°C, and the mean annual precipitation (MAP) varies between 223 mm and 1934 mm. Detailed 112 information in each mountain can be found in Table 1.

113 2.2 Soil sampling

114 Sampling was conducted at 166 sites across the 30 mountains. In each mountain, sites were selected 115 based on the altitude and dominant vegetation types. At each site, the geographic coordinate was recorded 116 using a GPS device (eTrex Venture, USA). Three replicate plots $(10 \text{ m} \times 10 \text{ m})$ were established, spaced 117 approximately 50 m apart to account for spatial variability. Soil profiles were manually excavated down 118 to the parent material in each plot. Soil horizons were delineated in the filed based on morphological 119 characteristics following the Chinese Soil Taxonomy (Chinese Soil Taxonomy Research Group, 2001; 120 Yang et al., 2023). Horizon boundaries were determined through visual and tactile assessments (e.g., 121 color, texture, consistency, moisture, and root distribution). Identified horizons were typically classified 122 as O (organic), A (surface mineral), and C (parent material) horizons. For each soil profile, the name,





- 123 code, depth range, and diagnostic features were recorded. Soil samples was collected separately from 124 each horizon in bottom-to-top sequence to avoid pollution. Samples were preserved in clean and sealed 125 polyethylene bags and transported to the laboratory. A total of 1314 samples were obtained, comprising 126 381 in O horizon, 481 in A horizon, and 452 in C horizon. The limited development of mountain soils 127 resulted in the absence of certain horizons in some profiles. All samples were air-dried, and visible roots, 128 litter, coarse fragments, and other debris were manually removed. The soils were then sieved through a 129 2 mm mesh prior to further analyses. Adjacent to each soil profile, bulk density (BD) was measured in 130 situ using stainless-steel cutting rings with different volumes depending on soil depth.
- 131 2.3 Soil physical and chemical analyses
- 132 Soil moisture content was determined by oven-drying at 105 °C. Soil pH was measured using a pH meter 133 (Mettler-Toledo FE28, Switzerland) after mixing soil samples with deionized water at a 1:2.5 ratio. Soil 134 organic carbon (SOC) concentration was determined by a CE400 elemental analyzer (Elementar vario 135 ISOTOPE cube, Germany), after removing carbonates with 5% HCl. Soil samples for element analysis 136 were digested with concentrated HNO₃, HF, and HClO₄ (Bing et al., 2022). The concentrations of major 137 elements (Al, Ba, Ca, Fe, K, Mg, Mn, Na, P, Sr, Ti, V, and Zn) in the digests were determined using an 138 inductively coupled plasma atomic emission spectrometry (ICP-AES, Optima 2000, USA), and the 139 concentrations of trace elements (Cd, Co, Cr, Cu, Mo, Ni, Pb, Sb, and Tl) were determined using an 140 inductively coupled plasma mass spectrometry (ICP-MS, Agilent 7700x, USA), with SPEXTM serving as 141 the standard solution. Quality control was ensured by analyzing replicates, blanks, and reference material 142 (BW07405, China). The recovery of the reference material was routinely within the range of 95-105%, 143 and the precision and accuracy of the analyses were < 5% (relative standard deviation).
- 144 **2.4 Environmental data extraction and calculation**

The MAT and MAP in the study area were collected from the WorldClim database (https://www.worldclim.org) with a resolution of 1 km. The aridity index (AI), calculated as the ratio of MAP to potential evapotranspiration (PET), was sourced from the Global Drought Index and Potential Evapotranspiration Database provided by the Plant Data Center of Chinese Academy of Sciences (doi.org/10.6084, CSTR:34735.11.PLANTDATA.0065, Zomer et al., 2022). The normalized difference vegetation index (NDVI) from 2001 to 2015 was derived from the Advanced Very High-Resolution Radiometer (AVHRR) dataset developed by the Global Inventory Modeling and Mapping Studies





- 152 (GIMMS) group (https://ecocast.arc.nasa.gov/data/pub/gimms/3g.v1/), with a resolution of 1/12°.
- 153 Atmospheric N deposition across China (1980-2015) were derived from Yu et al. (2019).

154 The chemical alteration index (CIA) was used to indicate the level of chemical weathering at each

155 site (Nesbitt & Young, 1982).

$$CIA = \frac{Al_2O_3}{(Al_2O_3 + Na_2O + K_2O + CaO^*)} \times 100$$

157 where the oxide is given in molar ratio, and CaO* refers to the amount of CaO incorporated in

158 silicate minerals.

156

159 3 Data description and evaluation

160 This dataset provides a comprehensive and continental-scale characterization of soil element composition 161 across China's mountains, spanning five major climatic zones and three pedogenetically distinct soil 162 horizons. The dataset includes quantitative measurements of 24 elements, thereby offering a holistic view 163 of soil geochemistry in mountain regions. The dataset integrates information from extensive field surveys, 164 laboratory analyses, high-resolution satellite-derived vegetation indices, and ancillary environmental 165 data compiled from national and global databases. To ensure data consistency and comparability across 166 sites, all data sources underwent rigorous harmonization procedures. In total, nine key environmental 167 drivers were compiled and categorized into four groups, including climate variables (e.g., MAT, MAP, 168 AI), vegetation characteristics (e.g., vegetation type, NDVI), basic soil properties (e.g., pH, moisture), 169 and atmospheric nitrogen deposition levels. These integrated variables enable robust evaluation of the 170 interactions between environmental factors and element distributions in soils. 171 3.1 Elemental concentrations across soil horizons

172Elemental concentrations in mountain soils exhibited a wide dynamic range, spanning over six orders of173magnitude. The highest mean concentration was observed for SOC (127,699 mg/kg), while the lowest174was for Tl (0.56 mg/kg, Fig. 2). The overall elemental mass ratio follows the pattern of C_{227038} : Al_{100998} :175 Fe_{46599} : K_{30204} : Ca_{19164} : Na_{14734} : Mg_{12201} : Ti_{5809} : P_{1373} : Mn_{1271} : Ba_{839} : Sr_{233} : Zn_{171} : V_{102} : $Cr_{87.1}$: $Pb_{71.1}$:176 $Ni_{30.3}$: $Cu_{28.4}$: $Co_{13.2}$: $Be_{3.31}$: $Sb_{2.92}$: $Mo_{2.04}$: $Cd_{1.18}$: Tl_1 , underscoring the compositional diversity shaped177by both natural and anthropogenic processes in mountainous environments.178Pronounced stratification of elemental concentrations was observed across the three soil horizons

179 (Fig. 3). The O horizon, enriched in organic matter, exhibited significantly higher concentrations of major





180 elements (e.g., carbon, Ca, and P), alongside elevated levels of several trace elements (e.g., Cd, Mn, Zn, 181 Pb, and Sb). This enrichment pattern is consistent with the accumulation of both biogenic nutrients and 182 atmospherically deposited pollutants. In particular, the elevated concentrations of Cd, Pb, and Sb in the 183 O horizon reflect anthropogenic inputs through long-range atmospheric transport and subsequent 184 deposition (Bing et al., 2019, 2021). These findings reinforce the role of the O horizon as a critical 185 biogeochemical interface where atmospheric inputs are retained and processed. In contrast, lithogenic 186 elements such as Al, Fe, K, Na, Mg, Ti, Sr, and V showed increasing concentrations with depth, 187 particularly in the A and C horizons (Fig. 3). This trend reflects the growing influence of parent material 188 composition and geochemical weathering processes with depth, where the contributions from organic 189 matter decline and pedogenic processes such as mineral dissolution and secondary mineral formation 190 dominate (Agnan et al., 2019; Kirkby et al., 2018; Woodruff et al., 2009). The systematic increase in 191 these elements suggests long-term pedogenic accumulation, consistent with the downward translocation 192 of weathering minerals and reduced biological cycling in subsurface horizons. Overall, this vertical 193 differentiation of elemental concentrations across soil horizons underscores the combined effects of 194 biological activity, atmospheric deposition, and geochemical weathering in shaping the composition and 195 distribution of soil elements in mountainous ecosystems.

196 3.2 Spatial distribution of soil elements across China's mountains

197 The spatial distribution of soil elements across China's mountains reveals complex patterns along 198 latitudinal, longitudinal, and altitudinal gradients (Fig. 4), indicating the integrated influence of climatic, 199 geological, and biological factors. Latitudinal trends in element concentrations were evident for several 200 elements. Notably, concentrations of K, Mn, Mo, Ba, and Sr increased consistently with latitude, 201 reflecting latitudinal variation in temperature-mediated weathering processes and element cycling (Moreno-Jiménez et al., 2022). In contrast, elements such as Ca, Mg, Na, Ni, Cu, Cd, Co, and Zn 202 203 exhibited unimodal distribution patterns, with peak concentrations in mid-latitude regions, particularly 204 at Mts. Suyukou, Wuyuezai, and Guandi. These non-linear patterns may arise from the combined 205 influence of regional precipitation regimes, variable soil development stages, and differential 206 atmospheric deposition (Bing et al., 2021; Luo et al., 2016a; Ren et al., 2019). Site-specific anomalies 207 were also observed, indicating the strong influence of localized geological conditions. For instance, the 208 high Ca concentrations at Mt. Gongga and elevated Ba and Sr concentrations at Mt. Dabie are attributed





- 209 to distinctive lithological features and mineralogical compositions at these locations, which can override
- 210 broader climatic or biogeographical trends (Yang et al., 2015; Zhi et al., 2004).
- 211 Elemental distributions along the longitudinal gradient also exhibited significant patterns. 212 Concentrations of Mn, Mo, and Zn increased from west to east across the study area, while Mg, Cr, and 213 V showed declining trends in the same direction (Fig. 4). Phosphorus exhibited a non-monotonic 214 longitudinal pattern, with concentrations decreasing initially and then increasing towards the eastern 215 regions. These longitudinal patterns may reflect transitions in geologic substrates, soil parent materials, 216 and regional differences in anthropogenic influence (Yang et al., 2022). Altitudinal variation in soil 217 elemental concentrations further highlights the influence of mountain-specific environmental gradients. 218 Concentrations of certain heavy metals (e.g., Cd, Cr, and V) increased significantly with altitude, while 219 Cu and Ni exhibited unimodal distribution patterns (Fig. 4). These altitudinal patterns are likely shaped 220 by altitude-driven shifts in local climatic conditions, vegetation type, and soil-forming processes, as well 221 as differences in parent material exposure and erosion dynamics (Moreno-Jiménez et al., 2022; Yang et 222 al., 2022).

223 3.3 Environmental drivers of soil element composition

224 The elemental composition of soils across China's mountains is significantly influenced by a combination 225 of climatic, edaphic, and biogeographic factors. Statistical analyses revealed that variables such as MAT, 226 MAP, AI, NDVI, CIA, soil pH, latitude, and soil moisture were significantly correlated with the 227 concentrations of multiple elements (p < 0.05, Fig. 5). Climate- and vegetation-related variables showed 228 negative correlations with SOC, P, Ca, Mg, Na, Mn, Ba, and Sr (p < 0.05). These relationships suggest 229 that cooler and drier sites, typically at higher altitudes or in arid regions, tend to accumulate higher 230 concentrations of these elements, due to reduced microbial decomposition, slower organic matter 231 turnover, and limited leaching under low-temperature and low-precipitation conditions (Moreno-Jiménez 232 et al., 2019, 2023). Conversely, warmer and more vegetated environments may enhance biological 233 cycling and leaching, leading to lower elemental retention in soils. Soil pH emerged as a key factor 234 influencing elemental availability and retention. Most elements (except SOC, Fe, Sb, Pb, Tl, and Ti) 235 exhibited positive correlations with pH (Fig. 5), indicating that higher pH conditions favor the retention 236 or reduce mobility of these elements. This pattern reflects the well-established role of pH in controlling 237 solubility, adsorption, and precipitation processes in the soil matrix (Barrow et al., 2023). Distinct vertical





238 patterns in elemental composition were observed across soil horizons, reflecting the transition from 239 biologically active surface layers to more geochemically stable subsurface zones (Bing et al., 2021; 240 Cronan et al., 2018; Jobbágy and Jackson, 2001). The O horizon exhibited elevated concentrations of 241 SOC, P, Ca, Sr, and Cd, corresponding to inputs from plant litter, root exudates, and atmospheric 242 deposition. The A horizon showed intermediate concentrations, indicating a mixing zone influenced by 243 both surface biological activity and subsurface geochemical processes. The C horizon was enriched in 244 lithogenic elements such as Ti, Al, Fe, and V, consistent with its proximity to parent material and 245 dominance of mineral weathering processes.

246 The explanatory power of environmental drivers varied by both elements and horizons (Fig. 6). In 247 the O horizon, climate and vegetation variables accounted for a substantial proportion of the variability 248 in Al, Be, and Cd, indicating a high sensitivity of surface layers to environmental inputs. In the A horizon, 249 the exceptionally high explanatory power observed for Fe (>200%) likely results from the combined 250 effects of redox conditions, clay mineral formation, and complexation with organic matter (Dong et al., 251 2023), although this warrants further investigation. In the C horizon, Na exhibited relatively high 252 explained variance, likely linked to its association with the weathering of Na-rich parent minerals. In 253 contrast, elements such as Sb, Tl, and V consistently showed low explanatory power across all horizons, 254 implying that their distributions are primarily controlled by lithological factors not captured by the 255 environmental factors in this study. The influence of environmental factors on SOC and other biologically 256 cycled elements decreased markedly with depth, consistent with the reduced biological activity and 257 increasing geogenic control in subsurface horizons (Jobbágy and Jackson, 2001; Sayer et al., 2020). In 258 contrast, pH and CIA became progressively more influential in the A and C horizons, highlighting the 259 increasing importance of geochemical stabilization and mineral transformation processes with depth.

260 4 Potential applications of the dataset

Soils are integral to global biogeochemical cycles, directly influencing ecosystem productivity, nutrient availability, and carbon storage. Understanding soil elemental composition and spatial variability is essential for assessing ecosystem function and predicting responses to environmental changes (Schimel et al., 2015). However, a persistent challenge in Earth system science has been the scarcity of highresolution, field-validated dataset, particularly in mountain regions where complex topography, diverse vegetation, and steep climate gradients complicate both empirical and model-based analyses (Luo et al.,





2016b; Tito et al., 2020; Todd-Brown et al., 2013). This comprehensive, multi-element soil dataset fills
a critical data gap by providing spatially explicit measurements of 24 elements across 30 mountain
regions in China, spanning five major climatic zones and three soil development horizons. It offers unique
opportunities for a wide range of applications in Earth system science.

271 Advancing biogeochemical modeling. The dataset is well-suited for parameterizing, calibrating, and validating a variety of biogeochemical models that simulate nutrient cycling, soil organic matter 272 273 dynamics, and element transport under changing environmental conditions. For instance, decomposition 274 models typically rely on the stoichiometry of carbon and nutrients, as well as interactions with 275 environmental drivers such as temperature, moisture, and pH (Fang et al., 2019; Feng and Zhu, 2021). 276 Our dataset provides detailed measurements of these variables across spatial and vertical gradients, 277 enabling more accurate estimation of model parameters that govern decomposition rates, mineralization, 278 and nutrient retention. Furthermore, the dataset facilitates model validation through direct comparison 279 between model outputs and observed elemental concentrations across diverse mountain ecosystems. This 280 is particularly valuable in regions where model performance has historically been limited by data scarcity. 281 By improving the representation of mountain soils in biogeochemical models, the dataset supports more 282 robust projections of nutrient cycling, soil carbon sequestration, and ecosystem feedbacks under various 283 climate and land-use change scenarios.

284 Enhancing soil quality assessment and ecosystem risk evaluation. Beyond modeling, the dataset 285 provides a valuable reference for assessing soil quality and ecological risks in mountain regions. 286 Measurements of SOC and macronutrient concentrations (e.g., P, K, and Ca) serve as indicators of soil 287 fertility and ecosystem productivity (Bauters et al., 2022; Cunha et al., 2022; Rizzo, et al., 2024). 288 Simultaneously, the inclusion of micronutrients and trace elements enables evaluation of potential 289 ecological hazards associated with heavy metal accumulation (Bing et al., 2021; Hou et al., 2025; Jin et 290 al., 2024). By capturing both beneficial and toxic elements, the dataset supports the identification of 291 nutrient-deficient or polluted soils, informing the development of region-specific soil quality benchmarks. 292 This is particularly relevant for mountainous ecosystems, where natural heterogeneity in soil properties 293 can lead to localized environmental vulnerabilities. Furthermore, the dataset offers a foundation for 294 integrated assessments of soil health and ecosystem services, including nutrient cycling efficiency, 295 pollutant buffering capacity, and biodiversity support.





296	Supporting scenario analysis and climate adaptation planning. By linking elemental concentrations
297	with environmental drivers (e.g., climate variables, vegetation indices, and soil chemical properties), this
298	dataset enables predictive analyses of how soil function may shift under future environmental conditions.
299	For instance, projections of climate change, altered precipitation regimes, or land-use intensification can
300	be used to model potential changes in soil nutrient availability, carbon storage, and heavy metal mobility
301	(Giovannelli et al. 2023; Ochoa-Hueso et al., 2023). These capabilities are essential for developing
302	adaptive land management strategies aimed at enhancing soil resilience and ecosystem sustainability in
303	vulnerable mountain regions.
304	Contributing to global soil and ecosystem monitoring initiatives. The dataset contributes to global
305	efforts in soil monitoring and environmental data sharing, such as those led by the Global Soil Partnership
306	(GSP), the International Soil Reference and Information Centre (ISRIC), and global biogeochemical
307	databases (e.g., SoilGrids, GEOTRACES) (Shi et al., 2025). Its high spatial resolution, standardized
308	sampling protocols, and comprehensive coverage make it a valuable input for global-scale assessments
309	of soil health, elemental cycling, and climate-soil-ecosystem interactions.
310	In summary, this dataset represents a significant advancement in the empirical foundation available
311	for Earth system research in mountainous environments. It offers broad applicability for improving
312	biogeochemical models, refining soil quality assessments, evaluating ecological risks, and informing
313	sustainable land management and climate adaptation strategies. By enabling deeper understanding of soil
314	processes across complex environmental gradients, this dataset contributes to ongoing efforts to conserve
315	and manage mountain ecosystems in the face of accelerating global change.
316	5 Data availability
317	The database is freely accessible at https://doi.org/10.11888/Terre.tpdc.302620 (Wu et al., 2025b). The
318	dataset provides comprehensive information for each sample, including mountain affiliation,
319	geographical coordinates, climatic characteristics, vegetation type, normalized difference vegetation

320 index, atmospheric nitrogen deposition rates, soil physicochemical properties, chemical weathering

321 indices, and concentrations of 24 soil elements.

322 Author contributions

BHJ, WuYH, and ZG designed the experiments. BHJ and LJ sampled the soils, BHJ, WangYH, YWZ,
and ZJ performed the analysis of soil samples. WYY, BHJ, and WangYH analyzed the data. WYY and





- 325 BHJ wrote the first draft of the manuscript. WYY, BHJ, WangYH, YWZ, ZJ, WuYH, LJ, and ZG revised
- and improved the manuscript. All authors read and approved the final manuscript.
- 327 **Competing interests**
- 328 The contact author has declared that none of the authors has any competing interests.
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559 Figure legends

- 560 Fig. 1 Geographic distribution of the 30 China's mountains. AL, Mt. Ailao; AS, Mt. Ao; BCW, Mt.
- 561 Baicaowa; CB, Mt. Changbai; DB, Mt. Dabie; DH, Mt. Dinghu; DX, Mt. Daxinganling; DYS, Mt.
- 562 Daiyun; FJS, Mt. Fanjing; GD, Mt. Guandi; GGS, Mt. Gongga; HS, Mt. Han; JF, Mt. Jifeng; JG, Mt.
- 563 Jiugong; JGS, Mt. Jinggang; LJ, Mt. Luoji; LG, Mt. Leigong; LJ, Mt. Luoji; ME, Mt. Maoer; NL, Mt.
- 564 Nanling; QF, Mt. Qingfengxia; QL, Mt. Qinling; SHB, Mt. Saihanba; SN, Mt. Shennongjia; SWDS, Mt.
- 565 Shiwandashan; SYK, Mt. Suyukou; TM, Mt. Tianmu; WGS, Mt. Wugong; WYZ, Mt. Wuyuezhai; WZS,
- 566 Mt. Wuzhi; XX, Mt. Xiaoxinganling.
- 567 Fig. 2 Frequency distribution of soil elements across the China's mountains. Red curve on each histogram
- represents the fitted normal distribution. The statistical parameters of the corresponding element are annotated in the upper left of each sub-figure.
- Fig. 3 Mean concentrations of 24 elements across different soil horizons. Lowercase letters indicate significant differences in each element among soil horizons (p < 0.05), and error bars represent the standard error.
- 573 Fig. 4 Spatial distribution of soil element concentrations across latitude, longitude, and altitude. The 574 colors of the dots represent different soil horizons. Solid red lines represent the fitting relationships of 575 elemental concentrations with latitude, longitude, and altitude (p < 0.05). R² less than 0.05 are not shown. 576 Fig. 5 The environmental factors influencing the elemental concentrations in the soils. (a) Redundancy 577 analysis (RDA) shows the relationships of soil element concentrations with environmental factors across 578 soil horizons. The inserted figure shows the distribution of samples along the axes. (b) Correlation 579 heatmap shows the correlation between soil element concentrations and environmental variables in each 580 soil horizon. The color and circle size represent the correlation coefficient, and * indicates statistical 581 significance (p < 0.05). SOC, soil organic carbon; MAP, mean annual precipitation; MAT, mean annual 582 temperature; DIN, dissolved inorganic nitrogen; NDVI, normalized difference vegetation index; BD, 583 bulk density; CIA, chemical index of alteration.
- Fig. 6 Explanation of elemental variation by environmental factors based on regression modelling. Columns with different colors represent different environmental variables. MAP, mean annual precipitation; MAT, mean annual temperature; DIN, dissolved inorganic nitrogen; NDVI, normalized difference vegetation index; CIA, chemical index of alteration.



588	Table1 The basic information across China's mountains $(n = 30)$
	Tuble1 The busie information deross enhalds information (in 50)

Climatic	timatic Mountain Latitude Longitude		Altitude	MAP	MAT			
zone	ID	n	(°N)	(°E)	(m)	(mm)	(°C)	Vegetation types
Cold-								Broadleaf forest, Coniferous-broadleaf
temperate	DX	133	49.25-53.45	122.34-124.28	360-1370	514	-3.1	mixed forest,
	BCW	36	40.82-40.83	117.60-117.61	1120-1710	539	6.0	Coniferous forest
	СВ	58	42.06-42.30	127.83-128.13	1025-2000	828	-1.0	Broadleaf forest, Coniferous forest
	HS	51	44.19-44.27	118.41-118.72	1100-1400	400	1.0	Broadleaf forest, Coniferous-broadleaf mixed forest,
Temperate	ME	12	45.39-45.39	127.68-127.68	400-400	599	2.7	Coniferous lorest Coniferous-broadleaf mixed forest Broadleaf forest
	SHB	36	42.33-42.47	117.29-117.51	1560-1870	452	1.0	Coniferous-broadleaf
	SYK	27	38.73-38.75	105.91-105.92	2030-2350	243	4.3	Coniferous forest
	XX	78	46.63-48.85	128.47-129.65	240-1420	639	0.2	Coniferous-broadleaf mixed forest
	AS	84	33.79-33.96	107.30-107.50	1260-3150	786	6.2	Broadleaf forest, Coniferous-broadleaf mixed forest, Coniferous forest Broadleaf forest,
	GD	42	37.89-37.90	111.43-111.44	1310-2200	5211	2.4	Coniferous-broadleaf mixed forest, Coniferous forest, Shrub
Warm-	JF	18	33.68-33.69	105.68-105.68	1785-1900	670	12.2	Coniferous forest
temperate								Broadleaf forest,
	QF	27	34.00-34.04	107.44-107.44	1530-2100	737	7.6	Coniferous-broadleaf mixed forest
	QL	66	34.02-34.16	107.61-107.80	870-2350	750	6.4	Broadleaf forest, Coniferous-broadleaf mixed forest, Coniferous forest
	WYZ	42	38.72-38.73	113.84-113.86	1210-1880	476	6.9	Broadleaf forest, Coniferous-broadleaf mixed forest,





								Coniferous forest, Shrub
	AL	45	24.50-24.54	100.99-101.03	2190-2655	971	14.1	Broadleaf forest, Coniferous-broadleaf mixed forest
	DB	45	31.09-31.10	115.77-115.81	400-1630	1562	11.4	Coniferous-broadleaf mixed forest, Coniferous forest
	DH	27	23.17-23.18	112.52-112.54	210-586	1788	21.1	Coniferous-broadleaf mixed forest
	DYS	30	25.64-25.65	118.22-118.22	1100-1510	1546	15.9	Coniferous-broadleaf mixed forest, Coniferous forest
	FJS	19	27.90-27.91	108.70-108.72	1482-2095	1346	11.6	Broadleaf forest
	GGS	66	29.54-29.60	101.96-102.07	2000-4225	967	6.6	Coniferous forest,
								Coniferous-broadleaf
	IG	27	29 38-29 40	114 58-114 62	580-1230	1635	13.8	mixed forest
	10		27100 27110	11.000 11.002		1000	10.0	Coniferous forest
Subtropical	JGS	30	26.51-26.63	114.11-114.17	925-1350	1727	14.0	Coniferous-broadleaf mixed forest
1								Broadleaf forest,
	LG	24	26.36-26.38	108.16-108.21	1223-2155	1294	12.6	Coniferous forest,
								Shrub
								Broadleaf forest,
	LJ	51	27.57-27.58	102.37-102.42	2200-3830	940	7.5	Coniferous forest
								Broadleaf forest,
		18	24.90-24.95	112.99-113.05	800-1545			Coniferous-broadleaf
	NL					1700	14.9	mixed forest,
								Coniferous forest
								Broadleaf forest,
	C 11	105	21.26.21.76			1074		Coniferous-broadleaf
	SN	105	31.36-31.76	110.15-110.55	1110-3090	1274	1.1	mixed forest,
								Coniferous forest
								Broadleaf forest,
	SWDS	45	21.88-21.90	107.90-107.92	393-708	1908	20.0	Coniferous-broadleaf
								mixed forest
	TM	21	30.36-30.58	119.43-119.73	790-1460	1382	14.0	Coniferous-broadleaf
								mixed forest
	WGS	27	27.46-27.46	114.16-114.17	778-1175	1767	13.6	Broadleaf forest
Tropical	WZS	24	18.89-18.90	109.69-109.71	862-1818	1694	20.8	Broadleaf forest





590	Table 2 The basic soil properties across the mountains $(n = 30)$
	Table 2 The basic son properties deross the mountains (n = 50)

Climatic zone	Mountain ID	Horizon	BD (g/cm ³)	Moisture (%)	рН	CIA
		0	0.11-0.21	70.0-182.2	3.37-6.73	53.40-67.58
Cold-temperate	DX	А	0.19-0.70	14.6-118.9	3.10-6.47	57.73-74.10
		С	0.50-1.39	9.1-44.6	3.70-6.44	57.6-74.04
		0	0.16-0.22	27.3-63.2	4.34-6.66	51.21-63.35
	BCW	А	0.24-0.59	25.4-37.5	4.88-6.54	53.82-62.86
		С	0.75-0.93	12.6-27.8	5.69-7.40	51.73-66.02
		Ο	0.11-0.15	103.8-224.5	4.83-6.32	46.31-62.36
	CB	А	0.18-0.31	86.2-152.0	4.34-6.44	49.90-62.77
		С	0.51-0.99	15.2-72.8	4.56-5.61	45.79-66.87
		О	0.21-0.26	19.2-48.3	5.51-6.82	53.33-58.89
	HS	А	0.32-0.70	18.3-27.0	5.16-6.62	53.11-58.54
		С	0.78-0.90	8.2-17.9	5.68-6.61	55.61-60.66
		О	0.16-0.16	59.1-59.1	4.72-5.88	58.78-62.16
Temperate	ME	А	0.38-0.38	42.0-42.0	4.49-5.24	58.08-61.88
		С	0.72-0.72	23.6-23.6	5.42-5.60	63.06-71.75
		О	0.16-0.23	25.0-66.9	5.70-7.08	58.82-62.28
	SHB	А	0.32-0.59	13.8-50.3	5.54-6.71	55.79-62.44
		С	0.84-0.92	7.5-22.0	5.36-6.44	57.28-61.71
		О	0.21-0.21	45.9-46.7	7.23-7.77	59.33-61.05
	SYK	А	0.25-0.36	34.9-45.6	7.28-7.92	58.76-59.95
		С	0.58-1.08	6.1-22.3	7.74-8.16	58.68-61.38
		Ο	0.11-0.19	35.6-109.6	4.64-6.35	58.92-63.85
	XX	А	0.30-0.54	30.5-60.1	4.11-6.09	57.78-66.81
		С	0.71-1.10	7.9-42.2	4.88-6.06	54.11-70.48
		О	0.19-0.56	61.6-237.3	3.70-6.32	59.18-68.21
	AS	А	0.24-0.63	45.9-204.9	4.05-6.63	57.08-70.85
		С	0.56-1.22	19.4-61.6	4.65-6.36	57.49-73.19
		О	0.13-0.20	131.2-152.1	5.41-6.67	58.72-63.30
	GD	А	0.25-0.59	61.1-127.3	5.63-6.72	57.39-59.84
		С	0.67-1.08	22.2-47.5	6.30-7.15	56.53-62.16
		Ο	0.22-1.15	24.0-94.4	4.18-6.31	66.41-69.56
	JF	А	0.21-0.55	40.0-180.0	4.27-5.40	66.71-70.03
Warm tamparata		С	0.73-1.02	10.6-35.8	5.48-6.30	68.69-76.31
warm-temperate		Ο	0.17-0.22	113.0-139.6	5.74-6.44	47.61-65.72
	QF	А	0.35-0.39	68.8-107.8	5.55-6.25	49.23-66.06
		С	0.59-0.88	30.2-47.3	5.96-6.41	49.89-65.65
		0	0.16-0.30	73.4-173.6	5.55-7.22	59.04-64.97
	QL	А	0.32-0.45	36.4-100.5	5.40-7.56	57.68-63.42
		С	0.73-1.20	8.1-37.5	5.60-8.11	50.17-64.64
		Ο	0.11-0.18	116.9-145.7	4.12-7.07	58.90-64.72
	WYZ	А	0.17-0.72	32.7-110.6	4.70-6.96	57.93-62.94
		С	0.77-0.96	15.3-22.3	5.81-6.33	60.68-66.56





		0	0.16-0.56	21.2-107.0	3.63-5.45	79.8-89.31
	AL	А	0.41-0.63	29.7-59.7	3.80-5.49	80.58-89.80
		С	0.69-0.92	24.8-45.7	4.55-5.53	79.87-91.00
		0	0.15-0.33	46.0-126.9	3.74-5.60	50.87-61.82
	DB	А	0.20-0.64	18.0-105.7	3.69-5.86	51.43-62.79
		С	0.82-1.03	8.5-26.2	4.66-5.90	51.45-74.48
		0	0.48-0.81	22.9-32.7	3.65-4.12	79.12-84.73
	DH	А	0.73-0.94	23.1-31.4	3.76-4.18	79.2-85.67
		С	0.88-1.15	15.5-17.9	3.97-4.32	79.95-85.76
	DVC	А	0.26-0.63	54.6-82.8	3.25-4.26	70.81-88.59
	DYS	С	0.53-0.97	32.4-49.4	3.58-4.31	77.30-89.18
		Ο	0.18-0.47	79.0-213.0	3.93-4.74	78.79-90.99
	FJS	А	0.24-0.40	66.3-197.0	4.06-4.73	79.95-92.02
		С	0.54-0.66	59.8-84.1	4.28-4.86	80.67-94.26
		0	0.15-0.24	58.5-73.1	3.39-6.76	52.55-63.91
	GGS	А	0.55-0.81	26.5-42.6	3.63-7.19	52.82-72.65
		С	1.11-1.30	13.4-29.4	4.28-8.02	48.12-75.17
		0	0.16-0.22	77.3-129.0	3.45-4.36	59.53-75.81
	JG	А	0.37-0.75	18.7-46.1	3.62-5.11	60.67-73.54
		С	0.68-0.89	14.6-19.1	4.36-5.22	55.5-74.38
	100	А	0.30-0.77	51.4-191.7	3.19-4.93	78.9-88.92
Subtropical	JGS	С	0.75-1.66	22.1-53.6	3.92-4.76	80.07-90.10
		0	0.43-0.48	77.5-78.9	4.11-4.67	76.37-86.52
	LG	А	0.49-0.74	18.8-70.3	4.26-4.71	79.38-87.29
		С	0.82-0.89	29.8-37.2	4.58-4.84	81.52-88.71
		0	0.18-0.36	74.5-173.1	3.59-3.91	54.95-69.98
	LJ	А	0.21-0.59	28.3-131.7	3.72-5.80	58.03-81.09
		С	0.85-1.02	19.5-40.1	3.83-6.34	59.00-82.45
		0	0.18-0.79	44.7-227.6	3.68-4.76	64.69-79.82
	NL	А	0.45-0.85	40.8-109.4	3.49-4.46	64.97-80.51
		С	0.83-1.12	14.7-43.5	4.12-4.61	66.8-83.91
		0	0.11-0.56	31.73-187.5	4.13-6.58	65.22-74.39
	SN	А	0.22-0.63	33.5-136.2	3.67-7.35	64.35-76.16
		С	0.55-0.96	20.5-61.3	4.16-6.78	68.13-81.39
		Ο	0.31-0.67	42.9-130.7	3.62-5.78	80.43-94.92
	SWDS	А	0.65-0.97	18.3-85.6	3.33-7.16	80.44-95.64
		С	0.84-1.52	17.3-41.0	3.69-4.28	80.69-95.81
		0	0.36-0.66	50.3-86.4	3.80-5.93	69.04-81.79
	TM	А	0.64-0.66	49.1-70.6	3.71-4.77	72.74-82.35
		С	0.92-0.92	29.9-29.9	4.58-4.66	74.97-75.77
	WCS	А	0.40-0.92	53.3-118.8	3.95-5.34	73.34-87.58
	w US	С	0.55-1.00	37.6-101.5	3.90-5.67	70.28-87.65
Tropical	WZS	А	0.71-0.86	49.7-73.7	3.52-3.96	84.89-92.80
ropical	WZS	С	1.04-1.38	25.2-32.2	3.98-4.42	85.40-93 40







Fig. 1 Geographic distribution of the 30 China's mountains. AL, Mt. Ailao; AS, Mt. Ao; BCW, Mt.
Baicaowa; CB, Mt. Changbai; DB, Mt. Dabie; DH, Mt. Dinghu; DX, Mt. Daxinganling; DYS, Mt.
Daiyun; FJS, Mt. Fanjing; GD, Mt. Guandi; GGS, Mt. Gongga; HS, Mt. Han; JF, Mt. Jifeng; JG, Mt.
Jiugong; JGS, Mt. Jinggang; LJ, Mt. Luoji; LG, Mt. Leigong; LJ, Mt. Luoji; ME, Mt. Maoer; NL, Mt.
Nanling; QF, Mt. Qingfengxia; QL, Mt. Qinling; SHB, Mt. Saihanba; SN, Mt. Shennongjia; SWDS, Mt.
Shiwandashan; SYK, Mt. Suyukou; TM, Mt. Tianmu; WGS, Mt. Wugong; WYZ, Mt. Wuyuezhai; WZS,
Mt. Wuzhi; XX, Mt. Xiaoxinganling.







Fig. 2 Frequency distribution of soil elements across the China's mountains. Red curve on each histogram represents the fitted normal distribution. The statistical parameters of the corresponding element are annotated in the upper left of each sub-figure.



















Fig. 4 Spatial distribution of soil element concentrations across latitude, longitude, and altitude. The
colors of the dots represent different soil horizons. Solid red lines represent the fitting relationships of
elemental concentrations with latitude, longitude, and altitude (p<0.05). R² less than 0.05 are not shown.







along the axes. (b) Correlation heatmap shows the correlation between soil element concentrations

and environmental variables in each soil horizon. The color and circle size represent the correlation

 $\label{eq:coefficient} 618 \qquad \text{coefficient, and * indicates statistical significance (p < 0.05). SOC, soil organic carbon; MAP, mean$

annual precipitation; MAT, mean annual temperature; DIN, dissolved inorganic nitrogen; NDVI,

620 normalized difference vegetation index; BD, bulk density; CIA, chemical index of alteration.







623 Columns with different colors represent different environmental variables. MAP, mean annual 624 precipitation; MAT, mean annual temperature; DIN, dissolved inorganic nitrogen; NDVI, 625 normalized difference vegetation index; CIA, chemical index of alteration.