

A global dataset of high-resolution CO₂ enhancements derived from OCO-3 measurements

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Abstract. We present a novel global dataset of CO_2 enhancements (ΔXCO_2) derived by fusing NASA's OCO-3 satellite and NOAA ground-based observations. CO_2 enhancements quantify the spatially resolved excess in atmospheric CO_2 concentrations arising from anthropogenic emissions, biospheric CO_2 exchanges, and atmospheric CO_2 transport. Leveraging decades of monthly CO_2 measurements from eight remote stations strictly selected from NOAA ESRL network, such as the Mauna Loa

- 5 station, we address the critical challenge of isolating localized CO₂ signals from background concentrations by developing a latitude-dependent global CO₂ baseline model that effectively captures spatial and seasonal variability in background CO₂. The developed baseline model demonstrates near-perfect hemispheric predictive accuracy (Northern: R^2 =0.988, RMSE=1.78 ppm; Southern: R^2 =0.995, RMSE=1.09 ppm). Spatially explicit ΔXCO_2 is then estimated by removing the column-corrected background CO₂ from co-located OCO-3 observations. Validations of the estimated ΔXCO_2 against tropospheric NO₂ (R^2 =0.896)
- 10 and prior in-situ urban CO₂ measurements, along with the dataset's high spatiotemporal resolution ($\sim 3 \text{ km}^2$), demonstrates its potential for tracking anthropogenic and biospheric CO₂ dynamics. Global Δ XCO₂ maps reveal mean CO₂ enhancements of 0.58 ± 1.81 ppm, with urban areas exhibiting 1.5-fold higher enhancements (1.43 ± 2.04 ppm). North Hemisphere land areas exhibits an approximately 81% higher Δ XCO₂ average (0.67 ± 1.98 ppm) compared to the South Hemisphere (0.37 ± 1.32 ppm), with urban enhancements amplifying this hemispheric contrast up to 95%. Comprising 54 million observations
- 15 across more than 200 countries, this open-access dataset provides an alternative metric for monitoring complex atmospheric CO₂ variability and actionable insights for regional climate policies, available at https://doi.org/10.5281/zenodo.15209825.

1 Introduction

Mapping global CO₂ enhancements with fine spatio-temporal resolution is essential for tracking anthropogenic and natural CO₂ sources, validating emission inventories and climate models, and assessing localized climate impacts, including effects on agricultural productivity. The CO₂ enhancement (Δ XCO₂) quantifies the combined influence of anthropogenic carbon

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emissions, the net CO₂ exchange between vegetation and the atmosphere, and large-scale atmospheric transport processes



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(Kiel et al., 2021; Mitchell et al., 2018; Reuter et al., 2019; Lei et al., 2022). A positive ΔXCO_2 signifies the net positive impacts of CO₂ sources and sinks with atmospheric transport processes on accumulation of atmospheric CO₂ concentration (Park et al., 2021; Kiel et al., 2021; Mitchell et al., 2018). ΔXCO_2 can become negative when enhanced terrestrial CO₂ uptake surpassing the amount emitted and atmospheric conditions such as prevailing wind result in a net decline in atmospheric CO₂ compared to long-term baseline. Accurate global ΔXCO_2 measurements are critical for understanding the current net

- contributions of local human and plant activities on atmospheric CO_2 increases. These fundamental measurements underpin the development of targeted strategies for future emission cuts and climate mitigation. Nevertheless, a comprehensive global dataset of CO_2 enhancements remains absent.
- 30 Carbon satellites have been reported as objective, independent data sources for monitoring spatiotemporal disparities of atmospheric CO_2 conditions (Pan et al., 2021; Nisbet and Weiss, 2010; Schwandner et al., 2017), by providing the top-down observations of column-averaged dry-air mole fraction of CO_2 (XCO₂; ppm). Satellite-derived CO_2 observations enhance data disclosure, transparency, and data equity especially in underdeveloped countries where data infrastructure is lacking and the accounting capacity of environmental departments is often weak. Moreover, the global coverage and high spatiotemporal
- 35 resolution of satellite instruments are key advantages that support the characterization of large-scale fine-grained atmospheric CO_2 levels. These capabilities, first, enable effective comparisons among multiple sites, cities, and countries; second, help refine the general patterns underlying local CO_2 variability, and ultimately, aid CO_2 mitigation efficacy by facilitating local governments making dynamic and targeted decisions.
- Mapping ΔXCO_2 based on satellite remote-sensing data remains challenging, although CO₂ satellites have been popular for 40 global atmospheric observations (Streets et al., 2013). The century-long persistence of CO₂ causes the immensely strong signal and notable spatial variability of background CO₂ concentration in the atmosphere, even in desert-like places (Hakkarainen et al., 2019). This accumulation masks the true CO₂ signals from local human and natural processes. Localized atmospheric CO₂ fluctuations are around two orders of magnitude smaller than the background CO₂ concentration (Canadell et al., 2023; Reuter et al., 2019). This substantial difference complicates the global differentiation of enhanced CO₂ signals from the accu-
- 45 mulative trend of background CO₂. Therefore, a key step in isolating localized ΔXCO_2 is deducting accurate fine-resolution background CO₂ concentration from the satellite CO₂ observations.

Satellite-driven ΔXCO_2 measurements are typically single site-specific, individual city-specific, or estimated from multiplesite gradients, rather than offering spatially-continuous coverage. Previous efforts to correct background CO₂ concentration have struggled with scale and seasonal sensitivity (Lindenmaier et al., 2014; Verhulst et al., 2017; Zeng et al., 2021; Miller et al.,

- 50 2020; Che et al., 2024; Kort et al., 2012; Schneising et al., 2013). These issues constrain the effectiveness in delivering largescale spatial analysis for fine-resolution ΔXCO_2 data. Site-specific ΔXCO_2 estimates are derived from the difference between satellite observation of atmospheric CO₂ and CO₂ records at ground-truth stations with minimal human and plant interference. These ground-truth stations include remote in-site stations from the Total Carbon Column Observing Network (Kiel et al., 2021) and mountaintop CO₂ observation sites in Salt Lake County, North America (Mitchell et al., 2018). City-specific ΔXCO_2
- estimates rely on the deviations of satellite observations in urban centers from the daily median (Hakkarainen et al., 2016; Park et al., 2021) or monthly median (Labzovskii et al., 2019) remote-sensing CO_2 observations in rural areas assuming that rural



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atmospheric CO₂ concentration should be considerably lower than that in urban cores (Wu et al., 2020; Reuter et al., 2019; Ye et al., 2020). The former site-specific approach cannot be effectively generalized for regional or global applications due to significant spatial heterogeneity in atmospheric conditions. The latter city-specific approach is sensitive to the seasonal timing of satellite overpasses, since seasonal biospheric-atmospheric CO₂ fluxes obscure the differentiation of urban-derived Δ XCO₂ signals from background CO₂ concentration in the rural areas. This limited generalizability can be in large part improved by

- signals from background CO₂ concentration in the rural areas. This limited generalizability can be in large part improved by leveraging the cooperative air sampling network of the National Oceanic and Atmospheric Administration (NOAA), which provide a range of remote marine stations for the developments of latitudinal references for background CO₂ concentration (Masarie and Tans, 1995; Tans et al., 1989). However, the continental-to-global spatial scales along multiple-site gradients of atmospheric background CO₂ (Mitchell et al., 2018) remain too coarse for precisely tracking background CO₂ dynamics
- in cities. This underscores the need for more spatially-explicit measurement techniques to better characterize subtle ΔXCO_2 variations within cities that contribute significant shares to global emissions (Duren and Miller, 2012).

To present both globally comprehensive and locally representative measurements of ΔXCO_2 , we leverage a NASA's new Orbiting Carbon Observatory 3 (OCO-3) satellite that offers the state-of-the-art highest resolution observations ($\approx 3 \text{ km}^2$)

covering the period from August 2019 to November 2023, and further remove atmospheric background CO_2 corrected from a novel global CO_2 baseline model based on ground-sourced CO_2 data from the Global Monitoring Laboratory (NOAA ESRL network). The satellite-derived dataset of global CO_2 enhancements enables objective, timely and spatially-explicit diagnosis of net impacts of CO_2 sources, sinks, and transport on atmospheric CO_2 increases, contributing to sub-city scale decision making on global net-zero strategies and climate actions.

75 2 Methods

This work develops a novel dataset of global CO₂ enhancements from 2019 to 2023 by integrating satellite-derived and groundsourced CO₂ observations. Fig. 1 demonstrates the dataset involved and the main workflows. The following three subsections elaborate on carbon satellite XCO₂ product (XCO₂), the global CO₂ baseline estimation (CO_{2b}), and the global CO₂ enhancements (Δ XCO₂), respectively.

80 2.1 Satellite-retrieved XCO₂ Observations

Satellite-derived XCO₂ observations (column-averaged dry air mole fraction of CO₂; ppm) are from the Orbiting Carbon Observatory 3 (OCO-3). This NASA satellite, launched in 2019, collects the global magnitude and distribution of atmospheric CO₂ concentrations with the highest spatial-temporal resolution to date, allowing it to track XCO₂ variations with a grid resolution of 2.2×1.6 km². Studies have proven that the OCO-3 is capable of detecting localized emission sources by giving

85 diurnal and geographically diverse XCO₂ observations (Kiel et al., 2021; Schwandner et al., 2017) and are less vulnerable to the impacts of small-scale atmospheric processes on the accuracy of local emission accounting (McKain et al., 2012). We use OCO-3 Level 2 bias-corrected XCO₂, version 10.4r data (OCO3_L2_Lite_FP), which is publicly available through the NASA Goddard Earth Science Data and Information Services Center (GES DISC) (http://disc.sci.gsfc.nasa.gov/). We filter the OCO-





Input Datasets







Table 1. The descriptive statistics of OCO-3 XCO₂ product.

Periods	Number of Observations	Mean (ppm)	1 st quantile (ppm)	99 th quantile (ppm)
2019-08 to 2019-12	3,807,526	408.49	402.39	413.67
2020-01 to 2020-12	15,045,735	411.52	405.17	416.70
2021-01 to 2021-12	13,526,529	414.17	408.36	419.57
2022-01 to 2022-12	12,383,251	416.44	409.86	421.94
2023-01 to 2023-11	9,685,130	418.49	412.58	424.02
Total	54,448,171	414.33	405.29	422.60

3 dataset using the 'Quality flag'. Quality flag 0 indicates high-quality remote-sensing data for scientific analysis. The full
window of August 2019 through November 2023 with 90,714,334 XCO₂ samples is covered, of which, a total of 54,448,171 terrestrial observations are used for developing global XCO₂ dataset (Table 1).

2.2 Global CO₂ Baseline Estimation

2.2.1 Selection of global CO₂ background stations

Global Monitoring Laboratory (GML) within the Earth System Research Laboratories (ESRL) of NOAA operates a wide net work of global ground-truth stations that accurately monitor long-term near-surface CO₂ levels. These stations are strategically deployed based on their geographical locations to represent background CO₂ concentrations.

The global CO_2 background stations are first required to be situated at considerable distances from urban infrastructures and anthropogenic activities to minimize human interference. Utilizing the global ESA Sentinel-2 10m land use/land cover map (Zanaga et al., 2022) to detect urban areas, we select monitoring stations that exhibit less than 2% built-up land cover

- 100 within a radius of 5 kilometers ($10 \times 10 \text{ km}^2$ area). To further minimize local 'contamination' from dense vegetation, we additionally select stations with an Enhanced Vegetation Index (EVI) of less than 0.3 within the predefined box size, based on MODIS Vegetation Index Products (Didan, 2021). Moreover, sea-level flask samples are excluded to minimize the influence of air-ocean exchanges. This approach, different from NOAA's marine station selection (Conway and others, 1994), focuses exclusively on terrestrial sampling stations and provides enhanced representations of terrestrial CO₂ patterns. Stations situated
- 105 within the Arctic and Antarctic circles, where carbon cycles exhibit unique characteristics (Bruhwiler et al., 2021), as well as those positioned downwind of emission sources, are similarly screened out. In addition, we conduct a visual inspection of the built-up surfaces and vegetation characteristics within a 5-km buffer zone around each monitoring station to avoid misguidance in station selection due to local heterogeneity. This examination is especially crucial for island stations, which often lack detailed land use and vegetation cover data. On this basis, we exclude stations with nearby vegetation or built-up
- 110 areas, even if these are minimal within the buffer zone. In accordance with baseline modelling requirements, we ensure that the selected background stations have CO_2 time series data for at least 10 years between 2000 and 2023.





Of the 88 active sites across 38 countries from the Global Monitoring Laboratory (NOAA ESRL network), a total of eight CO_2 background stations with wide latitudinal and longitudinal coverage (Fig. 2, Table 2) are chosen to model the global CO_2 baseline.



Fig. 2. The locations of eight CO₂ background stations. Station codes and names are highlighted in bold, including ICE: Storhofdi, Vestmannaeyjar; WLG: Mt.Waliguan; IZO: Izana, Tenerife, Canary Islands; ASK: Assekrem; MLO: Mauna Loa, Hawaii; NMB: Gobabeb; CRZ: Crozet Island; PSA: Palmer Station, Antarctica. Sources of Basemap: Esri, TomTom, Garmin, FAO, NOAA, USGS, © OpenStreetMap contributors, and the GIS User Community.

115 2.2.2 Development of a global CO₂ baseline model

We construct a global model to estimate spatio-temporal CO_2 baseline (CO_{2b}) using a simple and interpretable near-sinusoidal growth function (Eq.1). The mathematical form is an addition of a sinusoidal function and a linear function, both depending on a time parameter, m, denoting the number of months since January 2000. While the sinusoidal function simulates the wave-like periodicity and oscillation of seasonal CO_2 swings caused by seasonal changes in ecosystem photosynthesis and respiration

120 (Arrigo et al., 1987; Hall et al., 1975), the linear function simulates the long-term trend of atmospheric CO₂ concentration, which has been documented for decades (Groves et al., 1978). $\pi/6$ is the unit angle of one month in the sinusoidal function. *f* is the frequency of CO₂ seasonal variations. ϕ_h signifies the phase shift due to the inconsistent interhemispheric seasonality. ϵ_h denotes a residual term. *h* can be North or South Hemisphere.



Table 2. The location, percentages of built-up area, vegetation greenness, and satellite images of CO_2 background stations. The red trianglesare the locations of background stations. Source: Google Maps (© Google, accessed July 11, 2024)

Station	Location	Built-up Percentage (%)	Average EVI	Data Range	Satellite Images
ICE	63.400°North, 20.288°West	0.00	0.14	1992-2023	
WLG	36.288°North, 100.896°East	0.00	0.21	1990-2023	
IZO	28.309°North, 16.499°West	0.31	0.17	1991-2023	
ASK	23.263°North, 5.632°East	0.00	0.05	1995-2023	
MLO	19.536°North, 155.576°West	0.00	0.00	1969-2023	
NMB	23.580°South, 15.030°East	0.23	0.07	1997-2023	
CRZ	46.434°South, 51.848°East	0.02	0.21	1991-2023	
PSA	64.774°South, 64.053°West	0.00	NoData	1978-2023	A A A A A A A A A A A A A A A A A A A



Further, we allow parameters, β_{2,h}, β_{1,h}, and β_{0,h}, in Eq.1 to be linearly dependent on latitude, *l* (Eq.2), considering that
latitudinal gradient exerts considerable influence on human activities (Lei et al., 2021), naturally terrestrial biota, and air transmission through the troposphere (Denning et al., 1995) and thereby affecting the spatiotemporal distributions of CO₂ baseline levels (Conway and Tans, 1999; Shim et al., 2013). Specifically, studies reported a non-linear latitude dependence of seasonal CO₂ amplitude from the South Pole to the North Pole (Yun et al., 2022; Forkel et al., 2016; Heimann et al., 1998). The seasonal CO₂ amplitude varies slightly with latitude in the South Hemisphere whereas it shows a reinforced trend at high
latitudes in the North Hemisphere. The growth rate of atmospheric CO₂ is also sensitive to latitude (Taylor and Orr, 2000). β_{0,h}(*l*) represents the latitude-dependent prehistoric CO₂ concentration. In addition, Eq.3 ensures that baseline estimations in the North and South Hemispheres are continuous at the equator.

$$CO_{2b,h}(l,m) = \beta_{2,h}(l) \cdot \sin(\frac{\pi}{6} \cdot m \cdot f + \phi_h) + \beta_{1,h}(l) \cdot m + \beta_{0,h}(l) + \epsilon_h, \quad h \in \{south, north\}$$
(1)

$$|35 \quad \beta_{i,h}(l) = \beta_{1,i,h} \cdot l + \beta_{0,i,h} + \epsilon_{i,h}, \quad i \in \{0, 1, 2\}$$
(2)

subject to,
$$CO_{2b,south}(0,m) = CO_{2b,north}(0,m)$$
 (3)

The global CO₂ baseline model is constructed using 2,097 monthly observations collected between 2000 and 2023 from eight selected CO₂ background stations. This model has achieved near-perfect goodness-of-fit with R^2 =0.988 and RMSE=1.78 for North-Hemisphere curves, and R^2 =0.995 and RMSE=1.09 for South-Hemisphere curves. Fig. 3 illustrates time-series CO₂ observations and estimations at eight CO₂ background stations. For model simplicity, the student's t-test are applied to examine the statistical significance of all model parameters and only the significant ones (p-value < 0.05) remain in the final model. Eq.4 and 5 represent estimated CO₂ baseline models respectively for the North and South Hemispheres. These should be able to offer accurate estimates of CO₂ baseline levels at any latitude without being constrained by ground-truth station layout.

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$$CO_{2b,north} = -0.625 \cdot \sin(\frac{\pi}{6} \cdot m \cdot 1.002 + 0.846) + 0.129 \cdot l \cdot \sin(\frac{\pi}{6} \cdot m \cdot 1.002) + 0.186 \cdot m + 0.035 \cdot l + 366.240$$
 (4)

$$CO_{2b,south} = -0.625 \cdot \sin(\frac{\pi}{6} \cdot m \cdot 1.002 + 0.846) + (1.180e^{-4} \cdot l + 0.186) \cdot m + 0.027 \cdot l + 366.240$$
(5)

$$CO_{2b}(0,m) = -0.625 \cdot \sin(\frac{\pi}{6} \cdot m \cdot 1.002 + 0.846) + 0.186 \cdot m + 366.240$$
(6)

150 2.2.3 Sensitivity Analysis

Parameter sensitivity analysis is performed to examine the robustness of the developed global CO_2 baseline model over different criteria for background station selection. Multi-criteria combinations of the percentage of built-up areas, the EVI-based vegetation greenness, and buffer size of background stations are used for this purpose. By setting the thresholds of built-up percentages at 2% and 5%, the thresholds of EVI values at 0.2 and 0.3, and the thresholds of buffer radius at 2.5 km and 5.0

155 km, we compare the selected CO_2 background stations and their fitting performances. The background CO_2 estimations are highly stable and not sensitive to changes in background station selection criteria (Table 3).





R² = 0.983 RMSE = 2.09

 $R^2 = 0.990$

2016

RMSE = 1.47

2020

2020

RANN II

2024

2024

2024

2016

Station Code = WLG

2012 Year

Station Code = ASK

2012

Year

Station Code = NMB

2008

2008





2008

2012

Year

2016

2020

Fig. 3. The performance of the developed global CO₂ baseline model at eight CO₂ background stations.



(7)

Buffer Radius: 2.5 km				
Built-up Percentage (%)	Average EVI	Model Performance		
- D Ø	< 0.2	North Hemisphere Model: R ² =0.989; RMSE=1.67		
< 2%		South Hemisphere Model: R ² =0.994; RMSE=1.10		
< 5%	< 0.2	North Hemisphere Model: R ² =0.989; RMSE=1.67		
< 5%		South Hemisphere Model: R ² =0.994; RMSE=1.10		
~ 2%	< 0.3	North Hemisphere Model: R ² =0.988; RMSE=1.78		
< 2%		South Hemisphere Model: R ² =0.995; RMSE=1.09		
< 5%	< 0.3	North Hemisphere Model: R ² =0.986; RMSE=1.88		
		South Hemisphere Model: R ² =0.994; RMSE=1.10		
Buffer Radius: 5.0 km				
Built-up Percentage (%)	Average EVI	Model Performance		
- 001	< 0.2	North Hemisphere Model: R ² =0.989; RMSE=1.67		
< 2%		South Hemisphere Model: R ² =0.994; RMSE=1.10		
~ 50%	< 0.2	North Hemisphere Model: R ² =0.989; RMSE=1.67		
< 5%		South Hemisphere Model: R ² =0.994; RMSE=1.10		
~ 2%	< 0.3	North Hemisphere Model: R ² =0.988; RMSE=1.78		
 ∠ /0 		South Hemisphere Model: R ² =0.995; RMSE=1.09		
< 5%	< 0.3	North Hemisphere Model: R ² =0.988; RMSE=1.78		
5 /0	< 0.5	South Hemisphere Model: $P^2 = 0.005$: PMSE = 1.00		

Table 3. The sensitivity test of the estimated global CO₂ baseline model

2.3 A global dataset of CO₂ enhancements

The CO₂ enhancements, ΔXCO₂, is estimated as the difference between remote-sensing CO₂ observations (XCO₂) and a column-average concentration equivalent of the CO₂ baseline (XCO_{2b}) (Eq.7), considering the vertical profile of atmospheric
CO₂ by altitude (Bischof et al., 1980). We derived XCO_{2b} by applying a linear calibration function to CO_{2b}, based on 28,397 co-located observations from OCO-3 and CO₂ background stations (R²=0.909). The calibration function, with both its coefficients and constant terms linearly dependent on latitude due to the sensitivity of atmospheric vertical mixing and circulation to latitude changes (Monte-carlo and Sun, 1985; Huth et al., 2008), can scale ground-sourced CO₂ observations to satellite-based column average concentration.

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$$\Delta XCO_2 = XCO_2 - XCO_{2b}$$



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3 Dataset description

This study develops a cutting-edge dataset featuring instantaneous and highest-resolution ($\approx 3 \text{ km}^2$) estimates of global CO₂ enhancements, which contains 54,448,171 Δ XCO₂ observations across terrestrial regions (between approximately 52.88°S and 53.58°N) from August 2019 to November 2023. Of them, 1,231,856 Δ XCO₂ samples locate in urban areas, as defined by a published global urban extents map (Zhao et al., 2022). The dataset's high spatiotemporal resolution enables precise detection of short-term CO₂ enhancements driven by human activities and natural processes. For long-term evolution of atmospheric CO₂ burden, monthly or annual aggregations are recommended.

However, the spatio-temporal coverage of ΔXCO_2 is incomplete owing to OCO-3 observations loss attributed to cloud coverage and other data quality issues (Table. 1). Users are advised to examine the data coverage within their specific study

175 area. Efforts are underway to mitigate this problem leveraging spatio-temporal interpolation techniques (Wu et al., 2024). Ongoing improvements of dataset coverage are expected as OCO-3 observations accumulate. A future work will involve crossverifying background CO₂ estimates at a regional scale to ensure the consistency and accuracy of baseline models at different spatial resolutions.

The global dataset of CO₂ enhancements is stored in a NetCDF file with data attributes including latitude, longitude, time, satellite-derived CO₂, and CO₂ enhancements. At the time of writing this article, this dataset has been updated to November 2023. Future updates will also be available alongside with newly available OCO-3 data and NOAA ESRL station data.

4 Results and discussion

4.1 Dataset validation

4.1.1 Validation against tropospheric NO₂ measurements

- 185 The validation of this developed CO_2 enhancements dataset is performed by comparing ΔXCO_2 values in urban areas with co-located tropospheric NO₂ columns. This is a well-accepted approach because of the CO_2 -NO_x co-emission (Dou et al., 2023; Huo et al., 2022; Lei et al., 2022). Due to its shorter lifetime (hours), NO₂ is less affected by long-distance atmospheric transport, making it easier to be detected (Reuter et al., 2014; Richter et al., 2005). It means that the NO₂ column variation can trace the enhanced urban CO_2 compared to the global baseline due to human activities or natural processes (Wang et al.,
- 190 2020). This study uses the Sentinel-5P TROPOMI Tropospheric NO₂ L2 product (https://disc.gsfc.nasa.gov/), which provides the tropospheric NO₂ column measures at a spatial resolution of 3.5×5.5 km². Observations with a quality assurance (qa) value higher than 0.75 are utilized here to ensure trustworthiness. We geospatially match the hourly Δ XCO₂ samples and the co-located NO₂ observed throughout the three-hour window covering one hour before and one hour after the specified hour of the day. We also perform aggregation by averaging the Δ XCO₂ values within the NO₂ grids. Since the NO₂ data
- 195 is coarser than our dataset, we only use NO₂ grids that have at least three ΔXCO_2 observations to illustrate the relationship







Fig. 4. Validation of global CO₂ enhancements (Δ XCO₂). (a) Relationship between the binned Δ XCO₂ and co-located average NO₂ values in built-up areas worldwide. The error bar denotes the 95% Confidence Interval. (b) Compare the hourly Δ XCO₂ on three days in Los Angeles with a previous study (Kiel et al., 2021). The error bar denotes standard deviations.

between binned ΔXCO_2 and average NO₂ values. Using ΔXCO_2 samples from global built-up regions, we confirm a strong correlation between the proposed CO₂ enhancements and NO₂ values, with R²=0.896 (Fig. 4a).

4.1.2 Comparison with existing urban CO₂ enhancements

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We also compare ΔXCO_2 data proposed in this article with a previous urban CO₂ enhancements paper focusing on the Los Angeles megacity which measured background concentration using a remote surface station (Kiel et al., 2021). Kiel et al. (2021) reported intra-urban enhanced CO₂ relative to the background CO₂ for specific one hour on four days. To mitigate the wind impacts on locally atmospheric CO₂, only urban CO₂ observations from three of those days are compared (Fig. 4b). We find that the hourly ΔXCO_2 estimates provided by our datasets align closely with those reported in Kiel et al. (2021). The Mann-Whitney U test, with a p-value of 0.4, indicates insignificant difference between the ΔXCO_2 averages of two datasets. This suggests a great agreement between our CO₂ enhancements dataset and prior works.

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4.2 Spatial patterns of global CO₂ enhancements

Fig.5a illustrates the spatial distributions of global ΔXCO₂ measurements aggregated to 2°×2° grids. From 2019 to 2023, the global land exhibits ΔXCO₂ ranging from -49.86 ppm to 21.25 ppm with an average of 0.58 ppm and a standard deviation of 1.81 ppm (Fig. 5b). North Hemisphere shows an average ΔXCO₂ at 0.67 ppm (std=1.98), which is approximately 81%
210 higher than the average ΔXCO₂ of 0.37 ppm observed in the South Hemisphere (std=1.32) (Fig. 5b). We note that major

CO2 enhancements are concentrated in Eastern and South-Eastern Asia as well as equatorial tropical forests (e.g., Indonesia



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forest cover and tropical savannahs) (Fig. 5a). High ΔXCO_2 levels in Eastern and South-Eastern Asia are primarily driven by anthropogenic factors including accelerated socioeconomic development and population growth. In contrast, the elevated ΔXCO_2 near the equator can be attributed to biospheric factors. First, frequent fires in the tropical savannahs of the Sahel-Sudano-Guinean region emit substantial amounts of CO₂ into the atmosphere (Van Der Werf et al., 2017), Second, tropical woody vegetation increasingly functions as net carbon sources due to degradation and deforestation (Baccini et al., 2017).

With fine spatio-temporal resolution, the global dataset of CO₂ enhancements also provides detailed information for monitoring and characterizing ΔXCO_2 within urban areas. There are significant heterogeneity in urban ΔXCO_2 levels across six continents (Fig. 6a). North America has the highest average (mean = 1.76 ppm; std = 2.04 ppm), displaying relatively lower 220 ΔXCO_2 in the east than the west, because larger quantities of plants in the east absorb atmospheric CO₂ and attenuate the ΔXCO_2 levels. Asia follows, with an average of 1.72 ppm and a standard deviation of 2.08 ppm. By comparison, Oceania has the lowest ΔXCO_2 levels, averaging 0.59 ppm (std = 1.27 ppm) from 2019 to 2023.

The ΔXCO_2 values observed in global urban areas range from -16.89 ppm to 18.66 ppm with an average approximation of 1.43 ppm and a standard deviation of 2.04 ppm (Fig. 6b). The urban average ΔXCO_2 worldwide being roughly 1.5 times

225 higher than the global terrestrial mean highlights the significant influence of urban environments on the rise of atmospheric CO₂ levels. Urban areas in the North Hemisphere have an average ΔXCO_2 of 1.52 ppm, about 95% higher than the average level in the counterparts of South Hemisphere (Fig. 6c-d), owing to populated city distributions in the North. Fig. 7 exemplifies the spatial maps of ΔXCO_2 across six global cities. The spatially-explicit ΔXCO_2 values at sub-city scale have been clearly depicted and varying ΔXCO_2 levels between urban areas are also shown.

230 4.3 Temporal dynamics of global CO₂ enhancements

The temporal patterns of global CO₂ enhancements differ between the North and South Hemispheres, both in terms of changes in monthly ΔXCO₂ averages and in their seasonal variability. In the North Hemisphere, the monthly trend of ΔXCO₂ mean has a statistically significant increase (p-value<0.05 tested by a linear least-squares regression) from 2019 to 2023 with a slope of roughly 0.03 ppm per month (R²=0.31). Within North hemisphere urban areas, the monthly ΔXCO₂ mean increases approximately 0.04 ppm per month (R²=0.33) (Fig. 8a). No notable trend in the monthly ΔXCO₂ averages has been detected in the South Hemisphere and its urban areas (Fig. 8b). Moreover, the North Hemisphere, regardless of terrestrial land or urban coverage, experiences more pronounced seasonal fluctuations compared to the South Hemisphere which has lower terrestrial vegetation biomass.

Seasonal disparities in average ΔXCO_2 are evident (Fig. 8c), with the highest average (mean = 1.05 ppm) occurring in autumn in which terrestrial ecosystem increases net carbon release due to stronger total ecosystem respiration relative to gross primary productivity (Tang et al., 2022; Piao et al., 2008). In contrast, spring experiences the lowest average ΔXCO_2 (mean = 0.02 ppm) because of the onset of growing season and reduced human heating and cooling demands (Shen et al., 2014). Human cooling- and heating-caused CO₂ emissions explain the increases in averaged ΔXCO_2 during summer (mean = 0.67 ppm) and winter (mean = 0.40 ppm), respectively, compared to spring.







Fig. 5. The spatial map of CO₂ enhancements (Δ XCO₂) in global land areas at grids of 2° latitude by 2° longitude for illustration purpose (a). The box-plots display the Δ XCO₂ distributions at the original resolution (2.2×1.6 km²) across land areas globally, as well as within the North and South Hemisphere (b). The 'whiskers' of box-plots extend to 1.5 times the Interquartile Range (IQR) from Q1 and Q3, and outliers are hidden. The green triangle denotes the mean value and the orange line signifies the median value.

245 5 Conclusions

We generate a global dataset of CO_2 enhancements relative to the long-term CO_2 baseline, intending to identify ways that satellite-derived retrievals can guide local net-zero target and to foster international partnerships toward achieving carbon neutrality. Through synergistic integration of satellite-derived and ground-sourced CO_2 observations, this dataset provides 54,448,171 and 1,231,856 ΔXCO_2 records in global land areas and global urban areas over 2019-2023, respectively, with high

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spatial ($\approx 3 \text{ km}^2$) and temporal (instantaneous) resolutions. The spatiotemporal patterns also reveal that this CO₂ enhancements dataset is capable of capturing and characterizing the variability in net effects of carbon sources, sinks, and atmospheric transport on atmospheric CO₂ increases, across the globe and down to minimal intra-city scale.

The core strength of this dataset is to deliver spatially-explicit maps that enable instantaneous monitoring and ensure continuous, equitable tracking of CO_2 enhancements at multiple scales (across continents, countries, and cities). By identifying hot-spots and diagnosing temporal trends via a standardized metric, it equips international and domestic policy-makers with

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Fig. 6. The spatial maps of CO₂ enhancements (Δ XCO₂) throughout global urban areas ($2^{\circ} \times 2^{\circ}$) and six regions ($1^{\circ} \times 1^{\circ}$) (a). The Δ XCO₂ distributions within urban areas are presented for the global scale (b), the North Hemisphere (c), and the South Hemisphere (d). The 'whiskers' of box-plots extend to 1.5 times the Interquartile Range (IQR) from Q1 and Q3, and outliers are hidden. The green triangle denotes the mean value and the orange line signifies the median value.







Fig. 7. The spatial maps of CO₂ enhancements (ΔXCO_2) of six global cities aggregated to 0.02° resolution grids.

evidence-based insights. This will support the development of targeted and well-informed strategies for hierarchical carbon governance, and ultimately enhancing effective responses to climate risks.

6 Data availability

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The global dataset of CO_2 enhancements is freely available at https://doi.org/10.5281/zenodo.15209825 (Zhou et al., 2025). This dataset has been organized in NetCDF format under the WGS84 geographic coordinate system. It is compatible with Python, R, and free GIS software for reading and post-processing. As the satellite-derived and ground-sourced CO_2 observations update, the dataset will be continuously released aiming to consistently support in global carbon management and climate change mitigation.







Fig. 8. The monthly variations of CO₂ enhancements (Δ XCO₂) in the North Hemisphere (a) and South Hemisphere (b). The monthly mean Δ XCO₂ values exhibit statistically significant upward trend over land and urban areas in the North Hemisphere (p-value<0.05) whereas no significant trend is observed in the South Hemisphere. Global seasonal cycles of CO₂ enhancements are shown at (c). The 'whiskers' of box-plots extend to 1.5 times the Interquartile Range (IQR) from Q1 and Q3, and outliers are hidden. The green triangle denotes the mean value and the orange line signifies the median value.



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Author contributions. Y.Z. conceptualized the data product and supervised the project. P.F. and Y.Z. designed methodology, analyzed data,
 and wrote the manuscript. J.L. and Y.X. contributed to methodological design and manuscript editing. B.H. and C.W. contributed to conceptualization and manuscript editing.

Competing interests. The authors declare no competing interests.

Acknowledgements. This study is funded by National Key Research and Development Program of China (2022YFB3903704). J.L. acknowledges the funding from USMILE European Research Council (ERC CU18-3746). B.H. acknowledges the Senior Research Fellow Scheme from the Hong Kong Research Grants Council (SRFS2324-4H02).

https://doi.org/10.1017/9781009157896.007, 2023.



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