



- 1 Reconstructing Sea Level Variability at the Ieodo Ocean Research
- 2 Station (1993–2023) Using Artificial Intelligence, Machine Learning,
- 3 and Reanalysis Integration
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### Abstract

13 This study presents a comprehensive approach for reconstructing a high-quality, continuous 14 monthly sea level time series at the Ieodo Ocean Research Station (IORS) from 1993 to 2023 15 using advanced artificial intelligence (AI) and machine learning (ML) models. After applying quality control to the in-situ KIOST data, including inverse barometric effect correction, 3σ 16 17 filtering, and a 75% data coverage threshold, we validated trends using nearby PSMSL tide 18 gauges and four ocean reanalysis datasets (CMEMS, GLORYS, ORAS5, HYCOM). The trend 19 analysis showed a higher rate of sea level rise from in-situ data (4.94 mm/yr, Oct 2003-Dec 20 2023) compared to satellite and model-based estimates (e.g., CMEMS: 3.53 mm/yr, Jan 1993-21 Dec 2023), suggesting localized sea level rise in the East China Sea. Initial gap-filling used 22 statistical models such as harmonic regression and regression-based climatology. A blended 23 approach combining climatology and trend components achieved the best accuracy (RMSE 24 ~0.056 m,  $R^2 = 0.688$ ). We then implemented various AI/ML models through an Iterative Imputer framework. Ensemble models (e.g., XGBoost) performed perfectly after 2003 but did 25 26 not generalize well before 2004. Deep learning models like LSTM and GRU effectively captured seasonal and nonlinear patterns post-2003, with LSTM achieving RMSE = 0.023 m 27 28 and  $R^2 = 0.95$ . Time series models Prophet and SARIMA-SIN successfully reconstructed the 29 full time series, with SARIMA-SIN estimating the highest trend (5.61 mm/yr). Multiple linear 30 regression using reanalysis data served as a baseline, but AI/ML models outperformed it in both 31 accuracy and generalization. This study provides a reproducible, interpretable, and physically 32 consistent framework for reconstructing sea level variability in semi-enclosed coastal seas.

33

34 Keywords: Sea Level Variability, Ieodo Ocean Research Station, AI-based Imputation,





35 Timeseries Reconstruction, East China Sea



## 36 1. Introduction

37	Long-term sea level observations are critical for monitoring regional oceanographic and
38	climatic changes, particularly in coastal and marginal seas where variability is amplified by
39	topographic and atmospheric forcing (Hamlington et al., 2020; Cazenave and Moreira, 2022).
40	The Ieodo Ocean Research Station, located at the intersection of the East China Sea (ECS) and
41	the Southern East Sea (Sea of Japan), plays a pivotal role in observing regional sea level
42	variability and marine environmental conditions (Han, 2020; Byun et al., 2021). As a
43	strategically important in-situ platform, its various interval sea level records provide valuable
44	information for understanding the dynamics of the Kuroshio Current system, East Asian
45	monsoon variability, and climate change-driven sea level rise (Ha et al., 2019; Xu et al., 2015;
46	Chang and Oey, 2011).

However, long-term in-situ observations are often interrupted by equipment failure, maintenance issues, or extreme weather conditions such as typhoons (Adebisi et al., 2021). These disruptions lead to temporal gaps in the observational records, which hinder the detection of trends, reconstruction of seasonal cycles, and validation of satellite and model-based products (Beguería et al., 2019). In regions like the ECS, where strong seasonal and interannual signals are present, accurate and realistic imputation of missing data is essential for scientific and operational applications (Lin et al., 2020; Han, 2020).





54	Various methods have been developed to address missing data in oceanographic and climate
55	time series (Kolukula and Murty, 2025; Lee et al., 2022). Traditional approaches include linear
56	interpolation, monthly climatological averages, and harmonic regression models (Schlegel et
57	al., 2019; Risien et al., 2022; Okkaoğlu et al., 2020; Arguez and Applequist, 2013). More
58	recently, advanced statistical and machine learning techniques have been proposed for gap-
59	filling, including Gaussian process regression, Kalman filtering, and neural networks (Vance
60	et al., 2022; Wenzel and Schröter, 2010; Wang, 2023). While these methods offer improved
61	flexibility and accuracy, they often require dense observations or training data, which may not
62	be feasible in long-term sparse records (Lee et al., 2022; Sarafanov et al., 2022; Park et al.,
63	2023). Interpretable and statistically robust methods remain essential for operational and
64	historical datasets such as IORS (Han and Lim, 2024; Han, 2020).
65	This study focuses on imputing, filling, and predicting gaps in the monthly sea level data at the
66	IORS over 1993-2023. We evaluate and compare three regression-based approaches: (1)
67	harmonic regression with annual and semiannual cycles plus a linear trend (Okkaoğlu et al.,
68	2020), (2) a regression-based monthly climatology model with a trend using calendar year
69	month dummy variables (Hyndman and Athanasopoulos, 2018), and (3) a pure climatology-
70	plus-trend approach based on aggregated monthly means and linear fitting (Brunetti et al.,
71	2014). Furthermore, we propose a realistic blending method that optimally combines harmonic





72 and climatological components to minimize reconstruction errors.

73	Also, we implemented an ensemble of statistical and machine learning (ML) models to
74	reconstruct missing monthly values at the IORS (Tong and Li, 2025; Bochow et al., 2025).
75	These models were selected to span a broad range of algorithmic families, including ensemble
76	tree-based methods (e.g., XGBoost, Random Forest, LightGBM, AdaBoost) (Niazkar et al.,
77	2024; Wu et al., 2024; Gan et al., 2021; Xiao et al., 2019), regularized linear models (Lasso,
78	Ridge) (Pan et al., 2025), proximity-based models (K-Nearest Neighbors) (Latif et al., 2024),
79	and neural networks (LSTM, GRU) (Sun et al., 2020; Tumse and Alcansoy, 2025). Each model
80	was trained to impute missing values using only observed sea level data from the original
81	dataset, applying the IterativeImputer framework with consistent hyperparameters for
82	comparability (Ramirez et al., 2023). Complementing these were two timeseries specific
83	models: Facebook's Prophet (Elneel et al., 2024), which decomposes time series into seasonal
84	and trend components, and SARIMA (Sun et al., 2020), which captures both non-seasonal and
85	seasonal autocorrelation structures. All models were evaluated using root mean square error
86	(RMSE) and coefficient of determination (R <sup>2</sup> ), calculated against observed (non-missing)
87	values (Boursalie et al., 2022; Siddig et al., 2021).

88 To contextualize the reconstructed series, we further compared the IORS observations with 89 monthly sea level data from four global ocean reanalysis products—Copernicus Marine





- 90 Environment Monitoring Service (CMEMS), Global Ocean Physics Reanalysis System
- 91 (GLORYS), Ocean Reanalysis System 5 (ORAS5), and Hybrid Coordinate Ocean Model
- 92 (HYCOM)—over 1993–2023 (Han et al., 2024; Long et al., 2023; Jin et al., 2023b; Cummings
- and Smedstad, 2014). Linear trends were estimated and visualized to assess the consistency
- 94 and fidelity of in-situ observations relative to reanalysis-based records.
- 95 Finally, we extended the imputation framework to a multivariate setting by incorporating these
- 96 reanalysis datasets as auxiliary predictors, enabling more robust and physically informed gap-
- 97 filling. By leveraging both statistical and AI/ML-based techniques, this study provides a
- 98 transparent, reproducible framework for reconstructing realistic monthly sea level time series
- 99 in data-sparse coastal environments.







100

Figure 1. Map showing the locations of sea level observation stations used in this study,
including the Ieodo Ocean Research Station (IORS, black square) and seven tide gauge
stations (red circles) from the PSMSL and regional networks. The map spans the East
China Sea, Yellow Sea, East Sea (Sea of Japan), and the western Pacific Ocean. Major
geographic regions and oceanic features labelled, including Korea (KOR), China (CHN),
Japan (JPN), and Taiwan (TWN); KS and TAS are Korea and Taiwan Straits, respectively.





## 108 2. Data and Methods

### 109 2.1. Data and Preprocessing

- 110 This study utilizes sea level observations from the IORS, collected between October 2003 and
- 111 December 2023. The data, provided by the Korea Institute of Ocean Science and Technology
- 112 (KIOST), include measurements at 1-minute, 10-minute, and hourly intervals (Kiost, 2025).
- 113 Data gaps occur intermittently due to sensor malfunctions, scheduled maintenance, and severe
- 114 weather events.

To ensure physical consistency with satellite altimetry, numerical models, and reanalysis
products, all sea level values were corrected for the inverse barometric effect (IBE) in Eq. (1)
(Han et al., 2024). The correction was computed using collocated atmospheric pressure
measurements:

119 
$$\eta_{\text{corrected}} = \eta_{\text{raw}} + \frac{(P - P_0)}{\rho g}$$
 (1)

120 Where  $\eta_{raw}$  is the observed and uncorrected sea level (m), P is local atmospheric pressure (Pa),

- 121  $P_0$  is the standard atmospheric pressure (101300 Pa),  $\rho$  is the assumed seawater density (1025
- 122 kg/m<sup>3</sup>), and g is gravitational acceleration (9.81 m/s<sup>2</sup>).
- 123 A multi-tiered quality control protocol was implemented to construct reliable timeseries for
- subsequent analyses in Figures 2, 3 and 4:
- 125 Quality Flag Filtering: Only values flagged as "good" based on instrument diagnostics





126	and metadata were retained (Gouldman et al., 2017) (Figure 2).
127	- $3\sigma$ statistical Filtering: Outliers exceeding ±3 standard deviations from the local mean
128	were excluded (Oelsmann et al., 2021) (Figure 2).
129	- 75% Threshold Rule: Daily and monthly averages were retained only if at least 75%
130	of the corresponding observations passed the above filters (Epa, 2000) (Figures 3 and
131	4).
132	The resulting dataset comprises clean daily and monthly timeseries, exhibiting a clear seasonal
133	cycle, interannual variability, and several data gaps that require reconstruction. These
134	timeseries served as the foundation for further statistical modeling and machine learning-based
135	gap-filling.



Figure 2. Time series of inverse barometric effect (IBE)-corrected in-situ sea level
observations at the IORS from 1993 to 2023. Gray dots represent 10-minute interval data





- 139 flagged as good quality. Green dots indicate corrected 10-minute data within ±3σ. Red
- 140 dots denote daily means within  $\pm 3\sigma$ , and blue dots show monthly means within  $\pm 3\sigma$ . Data
- 141 gaps reflect periods of sensor malfunction or data loss.



142

Figure 3. Sea level time series at the IORS from 2003 to 2023, based on in-situ observations corrected for the IBE. Green dots represent 10-minute sea level data. Red dots show daily mean values calculated only when at least 75% of 10-minute observations per day are available and within ±3 standard deviations. Blue dots denote monthly mean sea level, computed from daily means that meet the same 75% availability and ±3 criteria.







148

149 Figure 4. Filtered daily and monthly sea level time series at the IORS from 2003 to 2023.

150 Red dots represent daily mean values calculated when at least 75% of 10-minute in-situ

- 151 observations are available and within ±3 standard deviations. Blue dots indicate
- 152 monthly mean values derived from daily means that also meet the 75% data availability
- 153 and  $\pm 3\sigma$  filtering criteria.

154

- 155 2.2. Sea Level Reconstruction Using Regression Techniques
- 156 We applied three regression-based approaches to reconstruct missing values and estimate long-
- 157 term sea level trends:

- 159 2.2.1. Harmonic Regression with Trend (HR)
- 160 This method models the sea level time series as a combination of a linear trend and periodic





- 161 seasonal components. Specifically, it includes annual and semiannual sine and cosine terms to
- 162 capture seasonal variability in Eq. (2) (Young et al., 1999):

163 
$$y_t = \beta_0 + \beta_1 t + \beta_2 \sin(\omega t) + \beta_3 \cos(\omega t) + \beta_4 \sin(2\omega t) + \beta_5 \cos(2\omega t) + \epsilon_t$$
(2)

- 164 where  $\omega = 2\pi/12$ , and t is time in months since January 1993. It captures both long-term and
- 165 seasonal variability and is especially suitable for periodic signals.

167 2.2.2. Regression-Based Monthly Climatology with Trend (RC)

168 This method models monthly sea level as a function of a linear time trend and monthly

- 169 categorical effects. Multiple linear regression uses dummy variables  $D_m$  (0 or 1) for calendar
- 170 months (February through December), with January as the reference month. This allows the
- 171 estimation of monthly climatology and trends simultaneously in Eq. (3). The model is
- 172 expressed as (Hyndman and Athanasopoulos, 2018):

173 
$$y_t = \beta_0 + \beta_1 t + \beta_2 D_2 + \beta_3 D_3 + \dots + \beta_{12} D_{12} + \epsilon_t$$

$$174 = \beta_0 + \beta_1 t + \sum_{m=2}^{12} \beta_m D_m + \epsilon_t$$
(3)

175 Where  $y_t$  is the sea level at time t,  $\beta_0$  is the intercept,  $\beta_1$  captures the long-term linear trend, 176 and  $\beta_m$  quantifies the deviation of each month m from January. This formulation allows for a 177 flexible representation of the seasonal cycle without assuming a sinusoidal structure and can 178 accommodate non-harmonic monthly anomalies while still estimating a long-term trend.





180	2.2.3. Pure Climatology with Trend (PC)
181	In this method, seasonal climatology is first computed by averaging the observed sea level for
182	each calendar month across all years (Jin et al., 2023a). A separate linear trend is then fitted to
183	the full-time series. The two components are summed to reconstruct the entire series in Eq. (4).
184	While simple and interpretable, this method assumes stationarity in the seasonal cycle and may
185	miss interannual variations:
186	$y_t = C_{m(t)} + (\alpha_0 + \alpha_1 t) \tag{4}$
187	where $C_{m(t)}$ is the climatological mean for the month of t. This assumes a stationary seasonal
188	cycle.
189	
190	2.2.4. Realistic Blending (RB)
191	To enhance the realism of the reconstructed sea level time series, we implemented a Realistic
192	Blending (RB) approach that combines the outputs of the Harmonic Regression with Trend
193	(HR) and the Regression-Based Monthly Climatology with Trend (RC) models in Eq. (5). The
194	final imputed value at each time step is computed as a weighted average of the two models
195	(Krinner et al., 2005):
196	$y_t^{RB} = \alpha y_t^{HB} + (1 - \alpha) y_t^{RC} $ <sup>(5)</sup>





197	where $\alpha \in [0,1]$ is an optimal blending weight selected to minimize the root mean square error
198	(RMSE) between the blended estimate and the observed values. This method leverages the
199	physical interpretability of harmonic modeling and the statistical flexibility of regression-based
200	climatology, aiming to strike a balance between robustness and fidelity to observed variability.
201	
202	2.3 Machine Learning and Statistical Models for Data Gap Imputation
203	To fill gaps in the monthly sea level record, we evaluate a range of machine learning and
204	statistical models that have been effective for time series gap-filling in environmental datasets.
205	Each model is described briefly below with supporting literature:
206	
206 207	2.3.1. Extreme Gradient Boosting (XGBoost)
206 207 208	<ul><li>2.3.1. Extreme Gradient Boosting (XGBoost)</li><li>A powerful tree-based ensemble method that uses gradient boosting and inherently handles</li></ul>
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<ul><li>206</li><li>207</li><li>208</li><li>209</li><li>210</li></ul>	<ul><li>2.3.1. Extreme Gradient Boosting (XGBoost)</li><li>A powerful tree-based ensemble method that uses gradient boosting and inherently handles</li><li>missing values by learning default split directions (Niazkar et al., 2024). XGBoost has</li><li>demonstrated high accuracy in gap-filling environmental time series (e.g., used to impute</li></ul>
<ul> <li>206</li> <li>207</li> <li>208</li> <li>209</li> <li>210</li> <li>211</li> </ul>	<ul> <li>2.3.1. Extreme Gradient Boosting (XGBoost)</li> <li>A powerful tree-based ensemble method that uses gradient boosting and inherently handles</li> <li>missing values by learning default split directions (Niazkar et al., 2024). XGBoost has</li> <li>demonstrated high accuracy in gap-filling environmental time series (e.g., used to impute</li> <li>significant gaps in aerosol optical depth data successfully) (Fan et al., 2020)</li> </ul>
<ul> <li>206</li> <li>207</li> <li>208</li> <li>209</li> <li>210</li> <li>211</li> <li>212</li> </ul>	<ul> <li>2.3.1. Extreme Gradient Boosting (XGBoost)</li> <li>A powerful tree-based ensemble method that uses gradient boosting and inherently handles missing values by learning default split directions (Niazkar et al., 2024). XGBoost has demonstrated high accuracy in gap-filling environmental time series (e.g., used to impute significant gaps in aerosol optical depth data successfully) (Fan et al., 2020)</li> </ul>
<ul> <li>206</li> <li>207</li> <li>208</li> <li>209</li> <li>210</li> <li>211</li> <li>212</li> <li>213</li> </ul>	<ul> <li>2.3.1. Extreme Gradient Boosting (XGBoost)</li> <li>A powerful tree-based ensemble method that uses gradient boosting and inherently handles missing values by learning default split directions (Niazkar et al., 2024). XGBoost has demonstrated high accuracy in gap-filling environmental time series (e.g., used to impute significant gaps in aerosol optical depth data successfully) (Fan et al., 2020)</li> <li>2.3.2. Gradient Boosting Regressor</li> </ul>





- 215 iteratively to minimize prediction error. GBR has achieved low interpolation errors in climate
- 216 data gap-filling, outperforming neural networks and multiple linear regression in one study of
- 217 meteorological time series (Otchere et al., 2022; Körner et al., 2018).
- 218
- 219 2.3.3. Random Forest (RF)

220 An ensemble of decision trees trained on bootstrapped samples, aggregating their results.

221 Random forests can capture nonlinear relationships and have robust performance for missing

data imputation (Wu et al., 2024; Tang and Ishwaran, 2017). In comparative experiments, tree-

- 223 based methods like RF rank among the top performers for reconstructing missing
- 224 environmental time series data (Mahabbati et al., 2021).

225

226 2.3.4. AdaBoost

An adaptive boosting algorithm that sequentially combines many weak learners (often shallow trees) to improve prediction accuracy. The AdaBoost framework, while originally devised for classification, has a regression variant that has been applied to time series gap-filling, for example, by ensembling multiple neural-network base learners to impute missing traffic flow data, yielding improved accuracy and stability (Xiao et al., 2019; Shang et al., 2024).





233 2.3.5. LightGBM

234	LightGBM is a gradient boosting machine algorithm optimized for computational efficiency
235	and scalability, employing histogram-based feature binning and leaf-wise tree growth. It has
236	demonstrated high predictive accuracy in forecasting tasks and is widely applied in
237	hydrological gap-filling studies(Wang et al., 2025). Its ability to efficiently process large
238	datasets and natively handle missing values makes it highly suitable for gap-filling in extended
239	climate time series, while also effectively capturing the nonlinear interactions between tidal
240	dynamics and river discharge (Gan et al., 2021).

241

242 2.3.6. CatBoost (approximated via Gradient Boosting)

CatBoost is a gradient-boosting algorithm specifically designed to handle categorical features effectively. It introduces innovative techniques such as ordered boosting and efficient encoding of categorical variables, which help in reducing overfitting and improving model performance (Hancock and Khoshgoftaar, 2020). In this study, due to library constraints, we approximate CatBoost's functionality using scikit-learn's Gradient Boosting Regressor. While this approximation does not fully replicate CatBoost's specialized handling of categorical data, it allows us to implement a gradient boosting approach within our existing framework.





251	2.3.7.	K-Nearest Neighbors
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252	A non-parametric imputation approach that fills a missing value based on the values of its
253	closest neighbors in feature space. KNN is simple yet effective for gap-filling; a study
254	comparing time series imputation methods found that KNN achieved the highest reconstruction
255	accuracy for missing data in several cases (Ahn et al., 2022). The method leverages spatial or
256	temporal similarity, which can be particularly useful when neighboring station data or nearby
257	time points correlate with the missing sea level values.

259 2.3.8. Lasso and Ridge Regression

260 These are regularized linear regression models that impose L1 (Lasso) and L2 (Ridge) penalties, 261 respectively, to prevent overfitting. By shrinking coefficients, they provide stable estimates that 262 can be used to predict and interpolate missing values from other correlated variables. Regularized regressions have been explored in gap-filling contexts to utilize correlations in 263 264 environmental data while avoiding multicollinearity issues (Wijesekara and Liyanage, 2023). For instance, an elastic net (a combination of Lasso and Ridge) approach outperformed basic 265 ARIMA-based methods in imputing large gaps of air quality data, highlighting the value of 266 267 such penalized linear models.





269 2.3.9. Decision Tree

270	A single decision tree can be used as a regression model to estimate missing values by learning
271	piecewise constant relationships in the data. Although decision trees alone are prone to higher
272	variance, they have been applied to gap-filling and can capture nonlinear dependencies in sea
273	level time series. In practice, tree-based algorithms (and their ensembles) have been found
274	superior to many conventional interpolation techniques for climate data gaps(Zhu et al., 2023).
275	Simpler decision trees may serve as interpretable benchmarks, while often forming the building
276	blocks of more complex ensemble methods used for imputation.
277	

### 278 2.3.10. Prophet

A forecasting model based on additive decompositions of trend, seasonality, and holidays (developed by Facebook). Prophet is robust to missing data and irregular timing – it does not require regularly spaced observations and can model gaps without explicit interpolation (Elneel et al., 2024). It has been applied in environmental time series forecasting (e.g. for air quality and water levels) and can be used to predict and back-fill missing monthly sea level values by leveraging seasonal patterns and trends in the data.

285

286 2.3.11. SARIMA-SIN (Seasonal ARIMA with Harmonic Regression Preprocessing)





287	To enhance the capacity of traditional SARIMA models in capturing periodic signals and long-
288	term trends in oceanographic time series, we implemented a hybrid approach known as
289	SARIMA-SIN. This method integrates harmonic regression and SARIMA modelling (Wang et
290	al., 2019). In the first stage, a deterministic seasonal component is modeled using a linear
291	combination of sine and cosine basis functions representing annual and semi-annual cycles.
292	This harmonic regression accounts for the primary seasonality, after which residuals are
293	computed by subtracting the fitted seasonal signal from the original time series. This
294	preprocessing alleviates the need for SARIMA to fit both trend and periodicity simultaneously,
295	thereby improving model stability and interpretability. Compared to conventional SARIMA
296	models, SARIMA-SIN reduces overfitting and better preserves long-term variability, making
297	it particularly suitable for sea level gap filling where both seasonal dynamics and secular trends
298	are essential.

300 2.3.12. Neural Networks (LSTM and GRU)

301 Recurrent neural networks, particularly Long Short-Term Memory (LSTM) and Gated 302 Recurrent Unit (GRU) architectures, are well-suited for sequential data and have been 303 successfully applied to gap-filling in complex time series. These networks maintain internal 304 memory states that allow them to learn long-term dependencies, which is advantageous for





- 305 inferring missing segments of a sea level record. Studies have shown that LSTM models can
- 306 achieve more accurate gap imputations than traditional methods for environmental sensor data
- 307 (Song et al., 2020), and specialized RNN variants (e.g., GRU-D) are designed to handle
- 308 sequences with missing values by learning decay rates for missing inputs (Che et al., 2018). In
- 309 practice, LSTM/GRU networks can leverage the temporal patterns in sea level data (and
- 310 potentially exogenous inputs) to predict missing monthly values with high fidelity.
- 311
- 312 2.4. Evaluation Metrics and Implementation
- 313 To assess the performance of each imputation model, we computed two widely adopted
- 314 evaluation metrics:
- 315 2.4.1. Root Mean Square Error (RMSE)
- 316 RMSE quantifies the average magnitude of reconstruction error by penalizing large deviations
- 317 more strongly in Eq. (6) (Montgomery et al., 2021). It is defined as

318 RMSE = 
$$\sqrt{\frac{1}{N} \sum_{i=1}^{N} (y_i - \hat{y}_i)^2}$$
 (6)

where  $y_i$  are the observed values and  $\hat{y}_i$  are the reconstructed values at non-missing time steps.

321 2.4.2. Coefficient of Determination (R<sup>2</sup>)

322 The coefficient of determination,  $R^2$ , quantifies the proportion of variance in the observed data

323 that is explained by the model in Eq. (7) (Montgomery et al., 2021). It is calculated as:





324 
$$R^{2} = 1 - \frac{\sum_{i=1}^{N} (y_{i} - \widehat{y_{i}})^{2}}{\sum_{i=1}^{N} (y_{i} - \overline{y_{i}})^{2}}$$
(7)

325 Where  $y_i$ ,  $\hat{y}_i$ ,  $\overline{y}_i$ , and *N* denote the observed values, the model-predicted values, the mean of

326 observed values, and the number of valid (non-missing) observations.

327 An  $R^2$  value of 1 indicates a perfect fit, meaning the model explains all the variance in the data.

328 A value of 0 implies the model performs no better than simply predicting the mean, and

329 negative values indicate that the model performs worse than the mean-based prediction. Thus,

- 330 higher  $R^2$  values reflect stronger agreement between the reconstructed and observed time
- 331 series.

332

#### 333 2.5. Comparison with Ocean Reanalysis Products

334	To contextualize in-situ observations, monthly sea level data from four global ocean reanalysis
335	products-CMEMS, GLORYS, ORAS5, and HYCOM-were compared with the KIOST in-
336	situ data at IORS over the period 1993-2023. All datasets were aggregated to monthly means
337	and aligned in time for consistency. Linear trends (expressed in mm/year) were estimated using
338	least squares regression and applied to each time series with NaN values masked. The trend
339	slope was computed in meters per month and converted to millimeters per year by multiplying
340	by 12,000. Each dataset was visualized on a same plot with annotated trend values to assess
341	agreement and long-term consistency between in-situ and reanalysis records. The comparison





- 342 allowed us to assess model fidelity and characterize interannual-to-decadal sea level variability
- 343 across observational and modeled sources.
- 344
- 345 2.6. AI/ML-Guided Multivariate Imputation Using Ocean Reanalysis Data
- 346 To improve the fidelity of gap-filling in the KIOST sea level time series, we implemented a
- 347 multivariate imputation framework that integrates auxiliary predictors from four ocean 348 reanalysis products: CMEMS, GLORYS, ORAS5, and HYCOM in Table 1. Specifically, we 349 applied IterativeImputer from the scikit-learn package in Python, using multiple linear 350 regression as the underlying estimator. In this configuration, the missing values in the KIOST
- 351 in-situ observations were modeled as a function of the corresponding values from the reanalysis
- 352 datasets, assuming linear dependence among variables.
- Each imputation process was iteratively updated over 20 cycles to ensure numerical convergence and stability of the imputed results. In parallel, univariate imputation was also conducted using standalone time series models—Prophet, SARIMA-SIN, LSTM, and GRU trained solely on the historical KIOST observations. These models capture temporal dynamics without external predictors.
- 358 Model outputs were evaluated based on the consistency of reconstructed trends, visual 359 coherence in the temporal structure, and alignment with physical expectations. The integration





- 360 of regression-based multivariate imputation and AI/ML models forms an ensemble-based
- 361 approach that leverages both data-driven learning and physically informed predictors for robust
- 362 reconstruction of missing values.
- 363

#### 364 Table 1. Overview of the datasets used in this study, including their temporal resolutions

365 and time periods.

Data Source	Temporal Resolution	Time Period
IORS	1-min, 10-min, hourly	Oct 2003 - Dec 2023
CMEMS	Daily, Monthly	Jan 1993 - Dec 2023
GLORYS	Monthly	Jan 1993 - Dec 2023
ORAS5	Monthly	Jan 1993 - Dec 2023
НУСОМ	Monthly	Jan 1994 - Dec 2023

366

## **367 3. Results**

368 3.1. Comparison with Other Data

369 A linear regression performed on the  $3\sigma$ -filtered, IBE-corrected monthly mean time series 370 revealed a significant upward trend of approximately 4.94 mm/yr from October 2003 to 371 December 2023 and 5.43 mm/yr from January 2004 to December 2023. This value is consistent 372 with other in-situ coastal tide gauge trends in the ECS and reflects regional steric and mass-





driven contributions to sea level rise in Figure 5 and Table 2.

374

- 375 Table 2. Monthly data availability for five PSMSL tide gauge stations from 1993 to 2023
- 376 (372 months total). The table shows the available and missing data points, the total
- 377 months considered, and the corresponding coverage percentage. These stations were
- 378 used to compare sea level variability and trends with the IORS in the Northwest Pacific
- 379 region.

	Available (Month)	Missing (Month)	Total (31 years)	Coverage (%)	Trend (mm/yr)
KANMEN	348	24	372	93.5	5.5
LUSI	291	81	372	78.2	6.2
SEOGWIPO	368	4	372	98.9	3.6
FUKUE	368	4	372	98.9	3.4
NAKANO SIMA	351	21	372	94.4	4.2
NASE III	352	20	372	94.6	2.6
NAHA	370	2	372	99.5	3.0





382 Figure 5. Monthly mean sea level time series from the IORS (bold black) and seven





383	PSMSL tide gauge stations (colored lines) from 1993 to 2023. PSMSL records are
384	vertically adjusted by subtracting 5.8 m to align with IORS. Linear trends are indicated
385	in the legend in millimeters per year (mm/yr), highlighting the long-term rise of sea level
386	across the region.
387	
388	3.2. Statistical Gap-Filling Models
389	To reconstruct missing values in the monthly sea level time series, five statistical gap-filling
390	models were evaluated, as shown in Figure 6. These models include:
391	- Harmonic Regression with Trend (HR): A model combining annual and semiannual
392	sinusoidal harmonics with a linear trend (red line).
393	- Regression-Based Monthly Climatology with Trend (RC): A linear regression using
394	monthly dummy variables to represent the seasonal cycle with an added trend (green
395	line).
396	- Pure Climatology with Trend (PC): A model combining the monthly climatological
397	mean and an original linear trend (cyan dashed line).
398	- Realistic Blending (RB, $\alpha y_t^{HB} + (1 - \alpha) y_t^{RC}$ when $\alpha = 0.00$ ): An optimally
399	weighted composite of HR and RC based on root-mean-square error (solid black line).
400	- Forced Blending (FB, $\alpha = 0.50$ ): A non-optimized blend of HR and RC using a fixed
401	mixing ratio (magenta dotted line).





Visual inspection reveals that the blended models, particularly the RB approach which is same to RC, closely align with the observed time series during periods of complete data. Among all candidates, the RB model ( $\alpha = 0.00$ ) demonstrated the best performance, achieving the lowest RMSE (0.0559 m) and the highest coefficient of determination (R<sup>2</sup> = 0.688) (Table 3). The RC model achieved identical metrics, followed by HR with RMSE = 0.0574 m and R<sup>2</sup> = 0.671. The PC underperformed relative to the others, with the highest RMSE (0.0633 m) and lowest R<sup>2</sup> (0.600), highlighting the limitations of neglecting local interannual variability.



Figure 6. Monthly sea level reconstruction at the IORS from 1993 to 2023 using five statistical gap-filling models. Blue dots indicate the original in-situ observations. Colored lines represent different model reconstructions: Harmonic Regression with Trend (HR, red), Regression-Based Monthly Climatology with Trend (RC, green dashed), Pure Climatology with Trend (PC, cyan dash-dot), Realistic Blending (RB,  $\alpha = 0.00$ , black dashed), and Forced Blending (FB,  $\alpha = 0.50$ , magenta dotted). The comparison highlights





- 416 differences in seasonal and long-term signal reconstruction, with the RB model providing
- 417 the most consistent alignment with observed values across the record.
- 418
- 419 Table 3. Performance evaluation of five statistical models for monthly sea level 420 reconstruction at the IORS (1993–2023). The table compares root mean square error 421 (RMSE) and coefficient of determination ( $R^2$ ) between observed and reconstructed sea 422 levels. The Realistic Blending model ( $\alpha = 0.00$ ) achieved the lowest RMSE and highest  $R^2$ , 423 matching the Pure Climatology with Trend model, indicating that harmonic components 424 did not improve reconstruction skill in this case.

Model	RMSE (m)	$\mathbb{R}^2$
Harmonic Regression with Trend (HR)	0.0574	0.671
Regression-Based Monthly Climatology		
with Trend (RC)	0.0559	0.688
Pure Climatology with Trend (PC)	0.0633	0.600
Realistic Blending (RB, a=0.00)	0.0559	0.688
Forced Blending (FB, α=0.50)	0.0563	0.684

#### 426

427 3.3. Artificial Intelligence and Machine Learning Imputation

428 To enhance the reconstruction accuracy of missing monthly sea level data at the IORS, a

429 comprehensive set of artificial intelligence (AI) and machine learning (ML) models was





430	implemented. These models were trained on the $3\sigma$ -filtered monthly KIOST time series with
431	synthetically introduced gaps to enable rigorous validation.
432	The methods included:
433	- Ensemble learning models: XGBoost, LightGBM, CatBoost, Gradient Boosting,
434	Random Forest, AdaBoost
435	- Statistical models: SARIMA with sinusoidal regressors (SARIMA-SIN) and Prophet
436	- Regularized regressions: Ridge and Lasso
437	- Neural networks: Long Short-Term Memory (LSTM) and Gated Recurrent Unit (GRU)
438	- Other models: K-Nearest Neighbors and Decision Trees
439	Figure 7 visualizes the full reconstructed time series from each model alongside the original
440	KIOST observations. Prophet (red solid line) and SARIMA-SIN (dark green dashed) are
441	visually emphasized due to their ability to reconstruct the full 1993-2023 time span,
442	consistently capturing both seasonal cycles and long-term trends. In contrast, XGBoost and
443	Random Forest models preserved the observed values and interpolated missing data, resulting
444	in perfect fit metrics ( $R^2 = 1$ ) but lacking predictive reconstruction before 2004 (Table 4).
445	As shown in Table 5, most machine learning models except Prophet, SARIMA-SIN, LSTM,
446	and GRU achieved perfect scores (MAE = 0, RMSE = 0, $R^2 = 1$ ) because they directly reused
447	observed values during imputation rather than making independent predictions, resulting in





artificially inflated performance metrics. Among the models that generated actual predictions,
LSTM achieved the lowest MAE (0.018 m) and RMSE (0.023 m) with a high $R^2$ of 0.95,
indicating its strength in capturing nonlinear temporal dynamics. GRU followed closely (MAE
= 0.030 m, RMSE = 0.037 m, $R^2$ = 0.87), balancing interpretability and predictive capability.
Prophet (MAE = 0.040 m, RMSE = 0.054 m, $R^2 = 0.71$ ) and SARIMA-SIN (MAE = 0.044 m,
RMSE = 0.059 m, $R^2$ = 0.66) exhibited slightly higher error values but demonstrated strong
consistency across the full time span (1993-2023), particularly in capturing seasonal and long-
term trends-even before the availability of observed data in 2004. Thus, while tree-based
ensemble models appear most accurate numerically, deep learning and time series models offer
true predictive value, especially in extrapolating missing or historical sea level variations.
Table 4. Evaluation metrics for AI/ML-based sea level reconstruction models at the IORS
from 1993 to 2023. Models were trained using the $3\sigma$ -filtered monthly KIOST data with
artificially introduced gaps. Performance was assessed based on mean absolute error
(MAE), root mean square error (RMSE), and coefficient of determination ( $R^2$ ), calculated
using the valid observed values (N). Most models, including XGBoost, LightGBM, and
CatBoost, produced near-perfect scores (MAE = 0, $RMSE = 0$ , $R^2 = 1$ ) because they simply

reused original values without generalizing to missing segments. In contrast, Prophet and
SARIMA-SIN provided full-series predictions, capturing both seasonal and long-term
variations. Deep learning models (LSTM, GRU) also exhibited strong predictive ability,

468 though their outputs were limited to periods with sufficient training data. Notably, many





### 469 models performed poorly before 2004, where fewer observations were available to guide

### 470 imputation.

Model	MAE (m)	RMSE (m)	R <sup>2</sup>	N
XGBoost	0	0	1	183
GradientBoosting	0	0	1	183
CatBoost	0	0	1	183
KNeighbors	0	0	1	183
Lasso	0	0	1	183
RandomForest	0	0	1	183
LightGBM	0	0	1	183
AdaBoost	0	0	1	183
Ridge	0	0	1	183
DecisionTree	0	0	1	183
Prophet	0.040	0.054	0.71	183
SARIMA_SIN	0.044	0.059	0.66	183
LSTM	0.018	0.023	0.95	183
GRU	0.030	0.037	0.87	183

471

472 Tree-based models outperformed deep learning in both accuracy and training stability.

473 However, LSTM showed strengths in capturing subtle nonlinear dependencies during seasonal

474 transitions.







475

476 Figure 7. Monthly sea level time series at the IORS from 1993 to 2023 reconstructed using 477 a suite of statistical and machine learning (AI/ML) models. Original observations (blue 478 dots) contain data gaps that were filled using ensemble methods (XGBoost, LightGBM, 479 CatBoost), regularized regressions (Ridge, Lasso), tree-based regressors (DecisionTree, 480 RandomForest, AdaBoost), neural networks (LSTM, GRU), and time series models 481 (Prophet, SARIMA-SIN). Prophet (red) and SARIMA-SIN (dark green dashed) are 482 visually emphasized for their capacity to capture both seasonal cycles and long-term 483 trends. The comparison illustrates model fidelity and divergence in reconstructing 484 historical sea level variability at IORS.

485

486 3.4. Comparison with Reanalysis and Satellite Data

487 The monthly sea level time series at the Ieodo Ocean Research Station (IORS) was compared 488 with satellite altimetry and four reanalysis products: CMEMS, GLORYS, ORAS5, and 489 HYCOM (Figure 8). The in-situ tide gauge data from KIOST exhibited a linear sea level rise





490	of 4.94 mm yr <sup>-1</sup> during the period from October 2003 to December 2023, with a slightly steeper
491	trend of 5.43 mm yr <sup><math>-1</math></sup> when computed from January 2004 onward (Table 5). These values
492	exceed the trends derived from satellite and model-based datasets, suggesting possible biases
493	due to local effects, datum inconsistencies, or limited assimilation in some models. Among the
494	reanalyses, CMEMS showed the closest agreement with the observed seasonal and interannual
495	variability, yielding a trend of 3.53 mm yr <sup>-1</sup> . GLORYS and ORAS5 also captured similar
496	variability, though with lower trends of 3.09 mm yr <sup>-1</sup> and 2.27 mm yr <sup>-1</sup> , respectively. In contrast,
497	HYCOM presented almost no net sea level rise (-0.09 mm yr <sup>-1</sup> ), likely due to limitations in
498	boundary conditions or lack of data assimilation in the East China Sea. Figure 8 clearly shows
499	consistent seasonal patterns across all datasets, though the amplitude and interannual signals
500	vary. The observed KIOST time series displayed greater interannual variability after 2003,
501	corresponding to the availability of continuous in-situ measurements. Cross-correlation
502	analysis further supports the consistency between CMEMS and other datasets (Table 6). The
503	correlation between CMEMS and KIOST reached 0.90, with similar values for GLORYS (0.93)
504	and ORAS5 (0.90). Correlation with HYCOM was slightly lower at 0.86, reflecting its weaker
505	agreement with observed variability. Separate correlations using only post-2003 KIOST
506	records confirmed similarly strong relationships. These results underscore the importance of
507	in-situ validation for semi-enclosed marginal seas like the East China Sea, where regional





- 508 processes and bathymetry can significantly affect long-term sea level trends and their
- 509 representation in models.
- 510
- Table 5. Linear sea level trends (mm yr<sup>-1</sup>) at the Ieodo Ocean Research Station (IORS) from multiple observational and reanalysis datasets. Trends are calculated using monthly data from tide gauge observations (KIOST), satellite altimetry (CMEMS), and ocean reanalysis/model outputs (GLORYS, ORAS5, HYCOM). The table includes trends for KIOST over two periods: October 2003–December 2023 and January 2004–December 2023, to assess sensitivity to the start date.

Dataset	Category	Period	Trend (mm yr <sup>-1</sup> )
In-situ	Tide Gauge (KIOST)	Oct 2003 - Dec 2023	4.94
In-situ	Tide Gauge (KIOST)	Jan 2004 - Dec 2023	5.43
CMEMS	Satellite Altimetry	Jan 1993 - Dec 2023	3.53
GLORYS	Ocean Reanalysis	Jan 1993 - Dec 2023	3.09
ORAS5	Ocean Reanalysis	Jan 1993 - Dec 2023	2.27
НҮСОМ	Ocean Model Output	Jan 1994 - Dec 2023	-0.09

#### 517

Table 6. Cross-correlation coefficients between CMEMS satellite altimetry and other
sea level datasets at the IORS (1993–2023). The comparison includes in-situ tide gauge
data (KIOST), ocean reanalysis products (GLORYS, ORAS5), and a numerical ocean





Data truca	Data trma	Cross-correlation
Data type	Data type	coefficient
CMEMS (1993–2023)	In-situ (2003–2023 and 2004–2023)	0.90
CMEMS (1993–2023)	GLORYS (1993–2023)	0.93
CMEMS (1993–2023)	ORAS5 (1993–2023)	0.90
CMEMS (1993–2023)	HYCOM (1994–2023)	0.86

## 521 model (HYCOM). All values are based on monthly mean anomalies.

522

523



Figure 8. Comparison of monthly sea level time series at the IORS from 1993 to 2023,
using in-situ observations (KIOST) and four ocean reanalysis products: CMEMS,
GLORYS, ORAS5, and HYCOM. Each dataset is plotted with its respective linear trend
estimated over the entire period. The observed sea level trend from KIOST (4.94 mm/yr)





530	is higher than those from CMEMS (3.53 mm/yr), GLORYS (3.09 mm/yr), ORAS5 (2.27
531	mm/yr), and HYCOM (-0.09 mm/yr), indicating significant discrepancies in modeled
532	versus observed trends at this location. Differences may reflect varying assimilation
533	schemes, vertical resolutions, or forcing mechanisms in each reanalysis system.
534	

535 3.5. Comparison with AI/ML-Based Reconstruction

536	Before implementing machine learning-based imputation methods, we applied multiple linear
537	regression using four reanalysis datasets (CMEMS, GLORYS, ORAS5, HYCOM) to evaluate
538	their explanatory power for observed sea level variability at the IORS. This regression served
539	as a preliminary benchmark to assess how well physically-based reanalysis products capture
540	regional sea level dynamics. Following this, a suite of AI/ML models was employed to
541	reconstruct the monthly sea level anomalies at the IORS from 1993 to 2023 (Fig. 9). Each
542	model's time series was compared with the original KIOST observations, and corresponding
543	linear trends (mm/year) were computed to assess long-term performance. The original KIOST
544	dataset exhibits a 4.94 mm/yr trend, serving as a reference benchmark. Among ensemble
545	learning methods, XGBoost (4.42 mm/yr), GradientBoosting (4.32 mm/yr), CatBoost (4.32
546	mm/yr), and RandomForest (4.26 mm/yr) closely reproduced the observed trend and temporal
547	variability. The Prophet model yielded a slightly higher trend estimate (4.51 mm/yr) and
548	consistently reconstructed the seasonal cycle and interannual variations. Its smooth,





549	interpretable structure made it effective for long-term monitoring, despite slightly larger
550	residuals than ensemble models. The SARIMA-SIN approach, which integrates harmonic
551	seasonal components with ARIMA modeling of residuals, produced the highest trend estimate
552	(5.61 mm/yr). While effectively modeling periodic signals, its slight overestimation may stem
553	from residual autocorrelation or the rigid seasonal structure embedded in the model. In contrast,
554	linear regularized models, such as Lasso (3.74 mm/yr) and Ridge (2.98 mm/yr), underestimated
555	the long-term trend and failed to capture higher-order seasonal and nonlinear dynamics. Neural
556	network-based deep learning models showed intermediate performance: LSTM (3.90 mm/yr)
557	and GRU (2.94 mm/yr) successfully captured high-frequency variability but smoothed long-
558	term trends. Their underestimation may reflect challenges in extrapolating temporal
559	dependencies across extended historical periods. Overall, ensemble models and Prophet
560	offered the most balanced performance regarding accuracy, trend reconstruction, and
561	robustness to missing data. SARIMA-SIN remains a promising alternative for seasonality-
562	focused applications. At the same time, deep learning methods may benefit from additional
563	architecture optimization and hyperparameter tuning to better preserve secular trends in semi-
564	enclosed regions like the East China Sea.









567 Figure 9. Monthly sea level time series at the IORS from 1993 to 2023 reconstructed using 568 various AI/ML models. The original KIOST observations (blue) are compared with 569 estimates from ensemble methods (XGBoost, GradientBoosting, CatBoost, 570 RandomForest, AdaBoost), regularized regression (Ridge, Lasso), decision trees, nearest 571 neighbors, Prophet, and deep learning (LSTM, GRU). A seasonal-harmonic SARIMA-572 SIN model is also included. Linear trends (in mm/yr) are indicated in the legend. The 573 comparison highlights each model's capability to capture seasonal and interannual sea 574 level variability.

575

## 576 4. Discussion

577 This study demonstrates a comprehensive approach to reconstructing monthly sea level 578 variability at the Ieodo Ocean Research Station (IORS) over the period 1993–2023 using a 579 suite of statistical, machine learning (ML), and artificial intelligence (AI) models. The





580	reconstruction framework was motivated by the need to address temporal discontinuities in						
581	long-term in-situ observations, which are often caused by instrument failure, extreme weather,						
582	or logistical constraints (Adebisi et al., 2021). These data gaps significantly hinder trend						
583	detection, seasonal analysis, and validation against satellite and reanalysis products in the East						
584	China Sea (ECS)-a region characterized by strong seasonal and interannual variability						
585	(Hamlington et al., 2020; Lin et al., 2020).						
586							
587	4.1. Data Sources and Preprocessing						
588	A diverse set of observational and reanalysis datasets was employed, spanning from high-						
589	resolution in-situ KIOST observations to global oceanographic products such as CMEMS,						
590	GLORYS, ORAS5, and HYCOM (Table 1). Preprocessing involved inverse barometric effect						
591	(IBE) correction, $3\sigma$ filtering, and 75% coverage thresholds to ensure consistency and quality						
592	in derived monthly means. These preprocessing steps align with established best practices for						
593	climate-quality oceanographic datasets (Cazenave and Moreira, 2022). Figures 1-3 illustrate						
594	the geographical coverage and temporal continuity achieved after these steps.						
595							
596	4.2. Regional Validation with PSMSL Tide Gauges						

597 To assess the validity of the reconstructed IORS series, comparisons were made with seven





598	nearby PSMSL tide gauge stations across the Northwest Pacific. The high correlation and						
599	consistent linear trends (2.6-6.2 mm/yr) observed in these stations (Table 2, Fig. 5) reinforce						
600	the regional representativeness of IORS data. This agreement supports the premise that well-						
601	maintained tide gauge records are invaluable for benchmarking reconstructions in semi-						
602	enclosed seas (Han, 2020).						
603							
604	4.3. Statistical Model Performance						
605	We evaluated five classical statistical gap-filling techniques, including harmonic regression and						
606	regression climatology. Realistic Blending (RB) and Regression Climatology (RC) achieved						
607	the best results, with RMSE of 0.0559 m and $R^2 = 0.688$ (Table 3; Fig. 6). These findings are						
608	consistent with prior literature that emphasizes the effectiveness of incorporating seasonal						
609	structure and linear trends in climatic time series (Okkaoğlu et al., 2020; Hyndman and						
610	Athanasopoulos, 2018).						
611							
612	4.4. AI/ML Model Evaluation						

Following statistical baseline evaluation, we implemented 14 AI/ML models using an Iterative
Imputer framework. Ensemble learners such as XGBoost and LightGBM achieved perfect
post-September 2003 metrics by reusing observed values, but lacked extrapolation capability





616	during earlier data gaps. In contrast, deep learning models like LSTM and GRU generated					
617	realistic, nonlinear predictions (LSTM: RMSE = $0.023$ m, $R^2 = 0.95$ ), although they					
618	underperformed in estimating long-term trends and were limited to the post-September 2003					
619	data range (Sun et al., 2020).					
620						
621	Time series-specific models Prophet and SARIMA-SIN produced full-period reconstructions					
622	with moderate accuracy ( $R^2 = 0.71$ and 0.66, respectively), but demonstrated superior ability					
623	to capture seasonal and long-term variability, making them suitable for regions with sparse data					
624	(Elneel et al., 2024; Tumse and Alcansoy, 2025). Their trend estimates (4.51-5.61 mm/yr)					
625	approximate observational trends better than most ML alternatives (Fig. 7).					
625 626	approximate observational trends better than most ML alternatives (Fig. 7).					
625 626 627	approximate observational trends better than most ML alternatives (Fig. 7). 4.5. Cross-Dataset Trend and Correlation Analysis					
<ul><li>625</li><li>626</li><li>627</li><li>628</li></ul>	<ul> <li>approximate observational trends better than most ML alternatives (Fig. 7).</li> <li>4.5. Cross-Dataset Trend and Correlation Analysis</li> <li>Trend comparisons across observational, satellite, and model-based datasets (Table 5; Fig. 8)</li> </ul>					
<ul> <li>625</li> <li>626</li> <li>627</li> <li>628</li> <li>629</li> </ul>	<ul> <li>approximate observational trends better than most ML alternatives (Fig. 7).</li> <li>4.5. Cross-Dataset Trend and Correlation Analysis</li> <li>Trend comparisons across observational, satellite, and model-based datasets (Table 5; Fig. 8)</li> <li>revealed that the IORS trends (4.94 mm/yr from October 2003 and 5.43 mm/yr from January</li> </ul>					
<ul> <li>625</li> <li>626</li> <li>627</li> <li>628</li> <li>629</li> <li>630</li> </ul>	<ul> <li>approximate observational trends better than most ML alternatives (Fig. 7).</li> <li>4.5. Cross-Dataset Trend and Correlation Analysis</li> <li>Trend comparisons across observational, satellite, and model-based datasets (Table 5; Fig. 8)</li> <li>revealed that the IORS trends (4.94 mm/yr from October 2003 and 5.43 mm/yr from January</li> <li>2004) are substantially higher than CMEMS (3.53 mm/yr) and GLORYS (3.09 mm/yr).</li> </ul>					
<ul> <li>625</li> <li>626</li> <li>627</li> <li>628</li> <li>629</li> <li>630</li> <li>631</li> </ul>	<ul> <li>approximate observational trends better than most ML alternatives (Fig. 7).</li> <li>4.5. Cross-Dataset Trend and Correlation Analysis</li> <li>Trend comparisons across observational, satellite, and model-based datasets (Table 5; Fig. 8)</li> <li>revealed that the IORS trends (4.94 mm/yr from October 2003 and 5.43 mm/yr from January</li> <li>2004) are substantially higher than CMEMS (3.53 mm/yr) and GLORYS (3.09 mm/yr).</li> <li>HYCOM, notably, displayed a negative trend (-0.09 mm/yr), likely due to boundary condition</li> </ul>					
<ul> <li>625</li> <li>626</li> <li>627</li> <li>628</li> <li>629</li> <li>630</li> <li>631</li> <li>632</li> </ul>	<ul> <li>approximate observational trends better than most ML alternatives (Fig. 7).</li> <li>4.5. Cross-Dataset Trend and Correlation Analysis</li> <li>Trend comparisons across observational, satellite, and model-based datasets (Table 5; Fig. 8)</li> <li>revealed that the IORS trends (4.94 mm/yr from October 2003 and 5.43 mm/yr from January</li> <li>2004) are substantially higher than CMEMS (3.53 mm/yr) and GLORYS (3.09 mm/yr).</li> <li>HYCOM, notably, displayed a negative trend (-0.09 mm/yr), likely due to boundary condition</li> <li>limitations such as excluding glaciers and land ice sheets melting (Han et al., 2024; Jin et al.,</li> </ul>					





634	strongly with in-situ (r = $0.90$ ), GLORYS (r = $0.93$ ), and ORAS5 (r = $0.90$ ), consistent with
635	prior findings on their reliability in coastal sea level analysis (Cummings and Smedstad, 2014).
636	
637	4.6. AI/ML Gap-Filling After Regression with Reanalysis
638	Prior to ML-based reconstruction, we conducted multiple linear regression using CMEMS,
639	GLORYS, ORAS5, and HYCOM to gauge their ability to explain IORS variability (Fig. 9).
640	The regression models underestimated the observed trend, underscoring their limitations in
641	data-sparse, dynamically complex environments like the ECS. Subsequent AI/ML
642	reconstructions-especially from Prophet, XGBoost, and SARIMA-SIN-outperformed these
643	baseline models in both fidelity and interpretability. SARIMA-SIN provided the highest trend
644	(5.61 mm/yr), while ensemble methods such as CatBoost and RandomForest approximated the
645	KIOST trend closely (4.3-4.4 mm/yr). Deep learning models like GRU underpredicted trends
646	(2.94 mm/yr), suggesting the need for further hyperparameter tuning or auxiliary input
647	inclusion.
648	

649 In summary, this study responds to the challenges outlined in the introduction-such as sparse long-term observational data and limitations of traditional gap-filling-by offering a 650 651 transparent, interpretable, and data-driven framework for sea level reconstruction. Through the





integration of AI/ML techniques and reanalysis validation, we provide new insights into local
sea level rise trends and offer a practical methodology for monitoring climate change effects in
marginal seas.

655

## 656 5. Conclusion

657 This study developed and validated a robust methodology for reconstructing the monthly sea 658 level record at the Ieodo Ocean Research Station (IORS) from 1993 to 2023 using a 659 combination of statistical techniques, machine learning models, and auxiliary reanalysis 660 datasets. Preprocessing steps, including inverse barometric effect (IBE) correction, 3s filtering, 661 and data coverage thresholds, allowed for the construction of reliable time series from highresolution in-situ observations. Comparison with seven nearby PSMSL tide gauges confirmed 662 663 the regional validity of the observed trends. The IORS sea level trends (4.94 mm/yr for 2003-664 2023 and 5.43 mm/vr for 2004–2023) were significantly higher than those from satellite and 665 reanalysis datasets, highlighting localized sea level rise. High correlations between CMEMS and other reanalysis products (r = 0.90) affirm their utility for sea level reconstruction, although 666 667 they underestimated trends compared to in-situ observations. Statistical models such as regression climatology and realistic blending provided accurate reconstructions, achieving the 668 best performance in RMSE and R<sup>2</sup> metrics. Among AI/ML models, ensemble learners (e.g., 669 670 XGBoost, RandomForest) achieved perfect reconstruction metrics after September 2003 but 671 failed to predict values in earlier periods. Deep learning models, particularly LSTM (RMSE = 672 0.023 m,  $R^2 = 0.95$ ), effectively modeled nonlinear and seasonal variability, though their ability 673 was restricted to after September 2003, and their trend estimation remained underestimated





674	compared to statistical benchmarks. Time series models, Prophet and SARIMA-SIN,
675	demonstrated the strongest capability for full-period imputation, delivering consistent seasonal
676	cycles and long-term trend estimates, with SARIMA-SIN reaching 5.61 mm/yr. Before
677	applying ML models, multiple linear regression using four reanalysis datasets was conducted
678	as a baseline. The AI/ML reconstructions outperformed reanalysis-based approaches in both
679	accuracy and consistency with the observed KIOST trend. Overall, this study provides a
680	scientifically grounded and computationally robust workflow for gap-filling, trend detection,
681	and sea level variability analysis. The framework offers a valuable tool for long-term sea level
682	monitoring, climate diagnostics, and policy planning in marginal seas and other observationally
683	limited regions.

684

# 685 6. Data availability

The original sea level dataset at the IORS is available at https://kors.kiost.ac.kr/eng/accessData and https://doi.org/10.17882/97666 (Kiost, 2025; Jeong et al., 2023). For updates to the datasets, please contact the data provider by writing to ycmin@kiost.ac.kr.

689

# 690 7. Author contribution

MH writing – original draft, visualization, data collection, investigation. HL writing – review
and editing, supervision.





## 694 8. Competing interests

695 The contact author has declared that none of the authors has any competing interests.

696

## 697 9. Acknowledgments

698	Oceanic	and	atmospheric	data	at	the	IORS
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