

CAMELS-COL: A Large-Sample Hydrometeorological Dataset for Colombia

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Abstract. Catchment Attributes and Meteorology for Large-Sample Studies (CAMELS-COL) is a large-sample hydrological 10 11 dataset for Colombia that integrates daily meteorological and hydrological time series with comprehensive catchment 12 attributes. The dataset comprises daily precipitation, evapotranspiration, and temperature from CHIRPS and MSWX satellite sources and streamflow records from 347 gauging stations covering the range 1981-2022. Additionally, CAMELS-COL 13 14 provides a wide range of catchment attributes, including physiographic characteristics, climatic indices, hydrological 15 signatures, land cover, geology, and soil properties, derived primarily from official governmental sources. CAMELS-COL 16 follows the standardized framework of previous CAMELS datasets, such as those developed for Brazil and Chile, to ensure 17 consistency with global hydrological datasets. By incorporating Colombian catchments across diverse hydroclimatic regions, including the Andean, Amazonian, and Caribbean basins, this dataset extends the CAMELS initiative into tropical 18 environments, offering a unique resource for hydrological research. The analysed basins show low flood susceptibility. An 19 analysis of the aridity index reveals that 74.7% of basins have a subhumid climate, 20.7% are semiarid, and only 4.6% are 20 classified as humid. Flow elasticity to precipitation is highest in the Amazon and Orinoco, highlighting their greater streamflow 21 22 sensitivity to rainfall changes. The Base Flow Index underscores groundwater's crucial role in stabilizing and regulating 23 surface water, particularly in forested basins. The dataset supports studies on hydrological processes, extreme hydroclimatic events, climate change impacts, and water resource management. Public availability encourages scientific collaboration and 24 25 facilitates the inclusion of Colombian catchments in continental and global hydrological analyses. The dataset is accessible at 26 https://doi.org/10.5281/zenodo.15554735 (Jimenez et al., 2025).



27 1. INTRODUCTION

One of the primary objectives of hydrology is to understand hydrological processes in sufficient detail so that hydrological models can accurately simulate a wide range of hydrological environments across different spatial and temporal scales and under diverse environmental conditions (Gupta et al., 2014). However, this cannot be achieved solely through intensive research in a few heavily instrumented basins. In this context, large-scale hydrological research supported by extensive datasets from multiple basins enables a broader, solid formulation and understanding of hydrological processes (Addor et al., 2020).

Currently, a variety of datasets compile different hydrometeorological and physiographic information from basins in various parts of the world. In this sense, Arsenault et al., (2020) compile in "The Hydrometeorological Sandbox - École de technologies supérieure – HYSETS" the hydrological, meteorological data, and physiographic information of more than 14,000 catchments located in high latitude regions covering 1950-2018. On the other hand, Kratzert et al. (2023) compiled and standardized climatological, hydrological, and physiographic information for over 6,830 basins as part of the Caravan initiative, resulting in an aggregated and standardized representation of the CAMELS (Catchment Attributes and Meteorology for Large-sample Studies) datasets.

40 The CAMELS project was designed to improve the representation of hydrological processes across large spatial domains, making it possible to simulate a variety of hydrological environments under different climatic and topographic conditions. 41 42 What differentiates CAMELS from other existing hydro-climatological datasets is its specific focus on catchment attributes, 43 which include not only meteorological data but also physiographic, hydrological, geological, soil, and land cover features that are critical for hydrological modeling. CAMELS was created for the United States CAMELS-US (Addor et al., 2017), and 44 later expanded to Chile CAMELS-CL (Alvarez-Garreton et al., 2018), Brazil CAMELS-BR (Chagas et al., 2020), United 45 Kingdom CAMELS-UK (Coxon et al., 2020), Australia CAMELS-AUS (Fowler et al., 2021), Switzerland CAMELS-CH 46 (Höge et al., 2023), Germany CAMELS – DE (Loritz et al., 2024), France CAMELS-FR (Delaigue et al., 2024), Denmark 47 48 CAMELS-DK (Liu et al., 2024), and India CAMELS-INDIA (Mangukiya et al., 2024).

49 In Colombia, most hydrometeorological stations are concentrated in the Magdalena region, a mountainous area within the 50 Andes Mountain range. Limited funding for the Institute of Hydrology, Meteorology, and Environmental Studies (IDEAM), 51 which is responsible for monitoring and researching the country's climate conditions, has restricted hydrometeorological 52 observation in regions such as the Amazon, Orinoco, and Pacific (Quesada and Avila-Diaz, 2023). On the other hand, as 53 highlighted by Avila-Diaz et al. (2024), the Water National Study (ENA)—one of the main governmental resources for 54 assessing the state of the country's water resources, updated biannually—has not incorporated updates on the country's water situation in its most recent editions. Moreover, these editions have neither included analyses related to the El Niño-Southern 55 56 Oscillation (ENSO) nor comprehensively addressed the impacts of global climate change in the country.

57 In this context, the CAMELS-CO dataset seeks to contribute to the understanding of hydrological processes across 58 Colombia. Colombia is the fourth largest country in South America (~1,141,748 km²) with a vast range of topographical and



59 climatological conditions, as the Pacific Ocean, the Andes Mountains, and the Amazon Rainforest influence it. To address the 60 challenges of understanding the hydrological dynamics in this topographically and climatologically diverse country, Furthermore, this work will contribute to the understanding of regional hydrology in South America. This paper introduces a 61 62 comprehensive and unique dataset covering 347 catchments across Colombia, with a higher concentration in the Andes 63 Mountain range. The dataset includes: i) Observed daily streamflow time series from 347 stream gauges; ii) Daily meteorological data derived from satellite observations (temperature and precipitation); iii) Daily meteorological data and its 64 65 derived products (e.g., potential evapotranspiration - ETo) iv) Climatic indices, v) Hydrological signatures, and iv) Catchments attributes, including land cover, geology, and soil characteristics, are calculated from official information provided by 66 67 governmental institutions. One of the main contributions of CAMELS-COL is the calculation of the Curve Number using the 68 Soil Conservation Service (SCS) method for each of the analyzed catchments, which is a parameter widely used in peak flow 69 estimation.

70 All data are provided in an accessible and standardized format, adhering to the workflow of previous CAMELS datasets, 71 which allows for direct comparison. By integrating hydrological data from multiple catchments into a single, ready-to-use 72 resource, this dataset streamlines data collection and preprocessing, significantly reducing time and effort. It supports 73 hydrological research in Colombia, enhances the benchmarking, calibration, and validation of hydrological models, and aids 74 in flood and drought monitoring. Additionally, it contributes to climate change adaptation planning at the departmental level 75 and enables global and regional hydroclimatic comparisons. The paper is structured as follows: Sect. 2 describes the study area and data collection. Sect. 3 details the derived catchment attributes, including physiographic, hydroclimatological, geological, 76 77 land cover, and soil characteristics, along with a discussion of their spatial distribution. Finally, Sect. 4 summarizes the main results of the paper. 78

79 2 Data provided by CAMELS-COL

The CAMELS-COL dataset provides a comprehensive collection of hydrometeorological, physiographic, climate indices, and hydrological signatures for Colombia's main catchments, following similar works in South America, such as CAMELS-BR (Chagas et al., 2020) and CAMELS-CL (Alvarez-Garreton et al., 2018).

83 The dataset includes catchment identification and localization. The physiographic characteristics encompass drainage 84 area, perimeter, Gravelius compactness index, form factor, main channel length, equivalent slope of the main channel, the 85 concentration time by four different methodologies (Kirpich, Ven te Chow, Johnstone, and Corps Engineers) and the mean curve number of Soil Conservation Service (SCS) method. Additionally, the dataset includes the percentage distribution of 86 87 land cover types from MapBiomas-Colombia (MapBiomas, 2022), as well as the lithological and soil characteristics of the 88 catchments, derived from the Geological Map of Colombia and the Soil Map, provided by the Colombian Geological Survey and the Agustín Codazzi Geographic Institute, respectively. In all cases, priority was given to official sources data, to ensure 89 90 the highest possible representativeness and reliability of the information included in CAMELS-COL.





The hydrometeorological data includes daily streamflow observations, precipitation, potential evapotranspiration, and maximum, mean, and minimum temperatures for each catchment. The climate indices encompass the aridity index, frequency, and duration of high and low precipitation days. The hydrological signatures include the 5th and 95th daily quantile flow, runoff ratio, streamflow elasticity concerning precipitation, slope of the flow duration curve, base flow index, and the frequency and duration of high and low flow days. Table 1 describes the information included in the dataset, and the following sections present a detailed explanation of the data source, and the methodology used to estimate each index.



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Table 1 Summary of the data provided by CAMELS-COL.

Attribute class	Attribute name	Description	Units	Data source
	gauge_id	Catchment identifier (IDEAM code of streamflow)	-	IDEAM
	gauge_department	Gauge streamflow department provided by IDEAM	-	IDEAM
Catchment	gauge_lat	Gauge streamflow latitude (WGS-1984 System)	°N	IDEAM
mormation	gauge_lon	Gauge streamflow longitude (WGS-1984 System)	°E	IDEAM
	gauge_elev	Gauge streamflow elevation	m.a.s.l.	IGAC DEM
	area	Catchment area	km²	IGAC DEM
	perimeter	Catchment perimeter	km	IGAC DEM
	gravelius_index	Gravelius compactness index of catchment	-	-
	factor_form	Form factor of catchment	-	-
	streng_chanel	Streng of the main chanel	km	IGAC DEM
Physiograpic	equi_slope	Equivalent slope of main chanel	m/m	Silveira (2005)
characteristics	cn_catchment	Mean curve number of Soil Conservation Service (SCS) method.	-	Wu et al., (2024)
	tc_kirpich	Concentration time by Kirpich Equation	hours	Silveira (2005)
	tc_chow	Concentration time by Ven te Chow Equation	hours	Silveira (2005)
	tc_Johnstone	Concentration time by Johnstone Equation	hours	Silveira (2005)
	tc_Engi_Corps	Concentration time by Corps Engineers Equation	hours	Silveira (2005)
	pr	Daily precipitation of catchment	mm/dia	CHIRPSV2
	poten_evapo	Daily potential evapotranspiration of catchment	mm/dia	MSWX
Hydrometeorological	t_max	Daily maximum temperature of catchment	°C	MSWX
data	t_mean	Daily mean temperature of catchment	°C	MSWX
	t_min	Daily minimum temperature of catchment	°C	MSWX
	streamflow	Mean daily streamflow of catchment	m ³ /s	IDEAM
	aridity	Aridity, computed as the ratio of mean potential evapotranspiration (ETo) to mean precipitation	-	Arora (2002)
	high_prec_freq	Frequency of high precipitation days (\geq 5 times the mean daily precipitation)	days yr-1	CHIRPS v2.0
Climatic indices.	high_prec_dur	Average duration of high precipitation events (number of consecutive days \geq 5 times the mean daily precipitation)	days	CHIRPS v2.0
	low_prec_freq	Frequency of dry days (< 1 mm day ⁻¹)	days yr-1	CHIRPS v2.0
	low_prec_dur	Average duration of dry precipitation events (number of consecutive dry days)	days	CHIRPS v2.0

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Attribute class	Attribute name	Description	Units	Data source
	q_mean	Mean daily discharge	mm day ⁻¹	
	runoff_ratio	Runoff ratio, computed as the ratio of mean daily discharge to mean daily precipitation	-	
	stream_elas	Streamflow precipitation elasticity (i.e., streamflow's sensitivity to precipitation changes at the annual timescale, using the mean daily discharge as reference). See equation 7 in Sankarasubramanian et al. (2001), with the last element being P/Q not Q/P (i.e as presented previously in CAMELS-BR)	-	
	slope_fdc	Slope of the flow duration curve between the log-transformed 33rd and 66th streamflow percentiles	-	
Hydrological signatures.	baseflow_index	Baseflow index, computed as the ratio of mean daily baseflow to mean daily discharge, with the hydrograph separation performed using the Ladson et al. (2013) digital filter	-	IDEAM
	Q5	5% flow quantile (low flow)	mm dav ⁻¹	
	Q95	95% flow quantile (high flow)	mm day ⁻¹	
	high_q_freq	Frequency of high-flow days (>9 times the median daily flow)	days yr ⁻¹	
	high_q_dur	Average duration of high-flow events (number of consecutive days > 9 times the median daily flow)	days	
	low_q_freq	Frequency of low-flow days (< 0.2 times the mean daily flow)	days vr ⁻¹	
	low_q_dur	Average duration of low-flow days (number of consecutive days < 0.2 times the mean daily flow)	days	
	forest_perc	Percentage of catchment covered by forest formation	%	
	nat_non_forest_form_perc	Percentage of catchment covered by natural non-forest formation	%	
Land cover	agricul_livestock_perc	Percentage of catchment covered by agriculture and livestock area	%	Mapbiomas
characteristics.	non_vegeted_perc	Percentage of catchment covered by non-vegetated area	%	(2022)
	water_bodies_perc	Percentage of catchment covered by water body	%	
	non_identi_land_perc	Percentage of catchment covered by non-identified land cover	%	
	unconso_depo_perc	Percentage of catchment covered by unconsolidated deposit	%	
	hypabyssal_rock_perc	Percentage of catchment covered by hypabyssal rock	%	
Castaria	metamor_rock_perc	Percentage of catchment covered by metamorphic rock	%	Cána at
characteristics.	plutonic_rock_perc	Percentage of catchment covered by plutonic rock	%	Gómez et al., (2023)
	sedimen_rock_perc	Percentage of catchment covered by sedimentary rock	%	, ()
	volcaniclastic_rock_perc	Percentage of catchment covered by volcaniclastic rock	%	
	volcanic_rock_perc	Percentage of catchment covered by volcanic rock	%	



Attribute class	Attribute name	Description	Units	Data source
	inceptisols_perc	Percentage of catchment covered by inceptisols	%	
	entisols_perc	Percentage of catchment covered by entisols	%	
	andisols_perc	Percentage of catchment covered by andisols	%	
	urban_perc	Percentage of catchment covered by urban areas	%	
	alfisols_perc	Percentage of catchment covered by alfisols	%	
	vertisols_perc	Percentage of catchment covered by vertisols	%	
	aridisols_perc	Percentage of catchment covered by aridisols	%	
C - 1	water_bodies_perc	Percentage of catchment covered by water bodies	%	
characteristics.	ultisols_perc	Percentage of catchment covered by ultisols	%	IGAC
	spodosols_perc	Percentage of catchment covered by spodosols	%	
	coal_mine_pit_perc	Percentage of catchment covered by coal mine pit	%	
	histosols_perc	Percentage of catchment covered by histosols	%	
	oxisols_perc	Percentage of catchment covered by oxisols	%	
	eroded_misce_perc	Percentage of catchment covered by eroded miscellaneous	%	
	rocky_misce_perc	Percentage of catchment covered by rocky miscellaneous	%	
	mollisols_perc	Percentage of catchment covered by mollisols	%	
	perpetual_snow_perc	Percentage of catchment covered by perpetual snow	%	



110 3 Methodology

111 **3.1 Study area**

The research focuses on Colombia, a country with diverse topography ranging from sea level to 5,703 meters above sea level (m.a.s.l) in the Andes Mountains. Colombia is segmented into five primary hydrological regions: the Amazon, Pacific, Orinoco, Caribbean, and Magdalena (Figure 1a). The Amazon contributes 37% of the country's surface streamflow, the Pacific 14%, the Orinoco 26%, the Caribbean 9%, and the Magdalena 14% (Bedoya, and Gonzales, 2021).

The Colombian climate exhibits substantial diversity and spatial variability. The warmest average monthly temperatures 116 are observed in the lowlands of the Magdalena and Orinoco regions, where temperatures exceed 24°C. In contrast, the coolest 117 118 temperatures are found in the mountainous areas, ranging from 8°C and 20°C Figure 1b). Rainfall patterns exhibit significant 119 variability: the northern Caribbean is the driest region, with an average annual rainfall of 500 mm/year, while the Pacific region 120 is the wettest, receiving over 11,000 mm/year in annual precipitation (Figure 1c). Annual evapotranspiration generally varies 121 from 1200 to 1400 mm/year, but it rises to 1400-1600 mm/year in the lowlands of the Magdalena region and the northern 122 Caribbean (Figure 1d). Regarding land cover, and considering the 2022 MapBiomas report, it is observed that most of 123 Colombia's land cover is mainly associated with forest formations, agricultural and livestock areas, and natural non-forest 124 formations (Figure 1e). According to the official geological map of Colombia (Gómez et al., 2023), the predominant 125 Colombian lithology is mainly associated with sedimentary rocks and unconsolidated deposits (Figure 1f).







precipitation in Colombia during 1981-2010 (mm/year) reported by the IDEAM. e) Colombian land cover identified in MapBiomas 2022. f) Colombia during 1981-2010 reported by the IDEAM . c) Mean annual precipitation in Colombia during 1981-2010 (mm/year) reported by the IDEAM. d) Mean annual Potential Evapotranspiration in Colombia during 1981-2010 reported by the IDEAM, and d) Mean annual Figure 1: Study Area. a) Colombian topography and streamflow stations considered in the study. b) Mean monthly temperature in Colombian Geology Map reported by the Geological Service of Colombia.





127 3.2 Data collection

128 3.2.1 Streamflow Data and Watershed Delineation

129 Daily Streamflow data is obtained from 921 gauging stations managed by **IDEAM** 130 (http://dhime.ideam.gov.co/atencionciudadano/) as of February 2023. Following the methodology adopted in CAMELS-CL, 131 CAMELS-IND, and CAMELS-DE, only streamflow stations with at least 10 years of records starting from 1980 are 132 considered. Additionally, to ensure the representativeness of the information included in CAMELS-COL, only basins with at 133 least 70% data availability are evaluated (see Figure 2).

Once the eligible streamflow stations are identified, the corresponding drainage areas are delineated using the Digital Elevation Model (DEM) of Colombia provided by the Instituto Geográfico Agustín Codazzi (Available at: Modelo Digital del Elevación. Colombia. STRM 30 Metros, <u>https://www.colombiaenmapas.gov.co</u>, accessed Jun 2024). The delineation process was carried out using the WhiteboxTools package (Lindsay, 2016), with streamflow stations serving as the outlet

138 points for each basin.

After delineating the drainage areas, their consistency was evaluated. After delineating the drainage areas, their consistency was assessed. Basins with inconsistent delineations were excluded, particularly in cases where inconsistencies were attributed to limitations of the Digital Elevation Model (DEM) used. Additionally, in line with the methodology used in CAMELS-DE, only basins containing at least one central pixel with available gridded weather data are considered. This decision ensures consistency with previous CAMELS-datasets, which included basins where gridded or satellite-based climate data could be reliably assigned within the basin boundaries.

After this evaluation. 347 streamflow stations are selected for analysis, with 22 located in the Amazon hydrological region, 31 in the Pacific, 57 in the Orinoco, 55 in the Caribbean, and 182 in the Magdalena basin (Figure 1a). Approximately 47% of the streamflow stations are located at elevations below 500 meters above sea level (m.a.s.l.), 21% between 500 and 1000 m.a.s.l., 13% between 1000 and 1500 m.a.s.l., 9% between 1500 and 2000 m.a.s.l., 5% between 2000 and 2500 m.a.s.l., and the remaining 5% above 2500 m.a.s.l. (Figure 1a).









152Figure 2 Number of stations with at least 10 to 30 years of daily streamflow records, categorized by different153thresholds of data availability (90%, 85%, 80%, and 70% percentage of days with streamflow observations).

154 3.2.2 Climatological Information

155 To analyze the hydroclimatological conditions of each basin, precipitation and maximum and minimum temperature data are obtained from satellite sources. The maximum (T_{max}) and minimum temperatures (T_{min}) are extracted from the MSWX 156 dataset. MSWX is a global gridded meteorological dataset with 3-hourly resolution at 0.1°×0.1° available at the 157 www.gloh2o.org/mswx by GloH2O (Araghi et al., 2023; Beck et al., 2022). MSWX consists of four sub-products: the historical 158 159 record from 1 January 1979 to approximately 5 days ago (MSWX-Past), the near real-time extension with a 3-hour latency 160 (MSWX-NRT), the medium-range forecast ensemble covering up to 10 days (MSWX-Mid), and the seasonal forecast 161 ensemble spanning up to 7 months (MSWX-Long) (Beck et al., 2022). At the same time, potential evapotranspiration (ETo) 162 is estimated using satellite-derived temperature data, applying the Hargreaves method (Hargreaves and Samani, 1985). The 163 calculation follows the equations outlined by Proutsos et al. (2024) as follows:

$$ETo = 0.0023 \cdot R_a \cdot (T + 17.8) \cdot (T_{max} - T_{min})^{0.5}$$
^[1]

where: *ETo* is the potential evapotranspiration (mm day-1), R_a is the extraterrestrial radiation (mm day-1) estimated from latitude and Julian day, calculated based on latitude and Julian day, following the methodology presented by the FAO (2006) in "*Evapotranspiración del Cultivo: Guías para la Determinación de los Requerimientos de Agua de los Cultivos (Estudios*



167 *FAO: Riego y Drenaje)*", *T*, T_{max} and T_{min} are the daily mean (calculated as the average of T_{max} and T_{min}), maximum, and 168 minimum air temperatures in °C respectively recovered from MSWX dataset.

Precipitation information is retrieved from the Climate Hazards Group InfraRed Precipitation with Station data - CHIRPS 169 (Available at: https://www.chc.ucsb.edu/data/chirps). It is a quasi-global (50°S-50°N) dataset with daily data available from 170 1981 to the present with a horizontal resolution of 5 km (Funk et al., 2015). The CHIRPS algorithm uses three data sources to 171 calculate precipitation: a) Rainfall data from CHPclim, b) satellite-based precipitation estimates using TIR (Thermal Infrared), 172 173 and c) in situ measurements from rain stations (Dinku et al., 2018). This dataset is selected for its effectiveness in accurately 174 representing precipitation patterns across Colombia (Romero-Hernández et al., 2024; Arregocés et al., 2023; Valencia et al., 2023; Cepeda and Cañon, 2022). The dataset presents daily data from January 1981 to December 2022 because it is the 175 176 common range between climatological information and streamflow series.

The average climatic conditions of each basin are estimated by calculating the mean of the pixel values located within its boundaries. No interpolation method was applied, as the goal was to use the original satellite data directly, preserving its resolution and spatial characteristics.

180 3.2.3 Geology

181 The geological characteristics of the study catchments are obtained from the "Geological Map of Colombia 2023, Scale 182 1:1.5M," prepared by the GSC (Gómez et al., 2023). This map was created from a raster with a spatial resolution of 30.0 m. The 193 geological units represented on the map are chronostratigraphic units grouped according to the age and lithology of 183 the materials. For the age, the units were classified into six groups (Cenozoic, Mesozoic, Paleozoic, Neoproterozoic, 184 Mesoproterozoic, Paleoproterozoic), with the reference being the 2022 International Chronostratigraphic Chart (Cohen et al., 185 2022). Regarding lithology, the units were divided into seven groups, distinguishing between rocks and deposits 186 187 (Unconsolidated deposit, Sedimentary rock, Volcaniclastic rock, Volcanic rock, Hypabyssal rock, Plutonic rock, Metamorphic 188 rock). The geological data of GCS can be accessed through their official website at 189 https://www2.sgc.gov.co/MGC/Paginas/mgc 1 5M2023.aspx. To facilitate the interpretation of the results, only the primary lithology classification is presented in the main text; this means Unconsolidated deposit, Sedimentary rock, Volcaniclastic 190 191 rock, Volcanic rock, Hypabyssal rock, Plutonic rock, Metamorphic rock. However, the feature "Geologic_characteristics" in 192 the dataset includes the percentage distribution of each of the 193 geological units in each catchment. The lithology for each 193 catchment is estimated by intersecting the geological dataset raster with the shape of the area.

194 3.2.4 Land Cover

The land cover of each catchment is estimated using data from MapBiomas Colombia for the year 2022 (MapBiomas,
2022). The classification maps, derived from Landsat mosaics, generate thematic land cover maps. These maps are part of the



MapBiomas initiative, which updates them whenever classification algorithms are refined. The MapBiomas Colombia dataset consists of a raster with a spatial resolution of 30.0 m. The dataset was accessed and downloaded in June 2023 from <u>https://colombia.mapbiomas.org/segunda-coleccion-de-mapbiomas-colombia</u>. It is important to note that, as of that date, the most recent dataset available was from 2022. Land cover estimation is performed by intersecting the MapBiomas raster with the shape of the area.

202 The use of MapBiomas Colombia in this study is justified by its high spatial resolution and standardized classification 203 methodology (MapBiomas, 2022). Additionally, it is important to highlight that no official governmental entity currently 204 provides an updated land cover dataset covering all of Colombia. In this context, MapBiomas Colombia stands out as a reliable 205 source, as it involves researchers and specialists in remote sensing, programming, ecosystems, and land use, ensuring a scientifically robust classification. Furthermore, its methodology enables comparisons with catchments in other countries, 206 207 which is particularly relevant given that this project spans multiple nations, including Brazil, Argentina, Peru, and Ecuador. The MapBiomas dataset classifies land cover into 20 types (Table 2), divided into five major groups: 1) Forest Formation, 2) 208 209 Natural Non-Forest Formation, 3) Agricultural and Livestock Areas, 4) Non-Vegetated Areas, and 5) Water Bodies. To 210 facilitate the interpretation of the results, only the five major classifications are presented in the main text. However, the feature 211 "Land cover characteristics" in the dataset includes the percentage distribution of each land cover category considered in 212 MapBiomas.

Table 2 Land Cover classification in MapBiomas Colombia 2022 - Collection 2





16	Mining						
17	Other non-vegetated area						
Water body							
18	River, lake or ocean						
19	Aquaculture						
20	Glacier						
Not observed							

214 3.2.5 Soil

The basins' soils are characterized using data from the "Soil Maps of the Colombian Territory—scale 1:100,000," published on the Open Data portal by the Agustín Codazzi Geographic Institute (IGAC), accessible at <u>https://geoportal.igac.gov.co/contenido/datos-abiertos-agrologia</u>. This source is selected for its official status and accurate representation of Colombian territory.

According to IGAC, these maps display the spatial distribution of soil characteristics, determined through general surveys conducted in each department. The maps provide descriptions and interpretations of soil-forming environments, the physical, chemical, mineralogical, and morphological properties of the soil, its taxonomy, drainage capacity, and spatial distribution. The soil types featured in the Soil Maps of the Colombian Territory include: Alfisols, Andisols, Aridisols, Coal Mine Pit, Entisols, Eroded Miscellaneous, Histosols, Inceptisols, Mollisols, Oxisols, Perpetual Snow, Rocky Miscellaneous, Spodosols, Ultisols, Urban areas, Vertisols, Water Bodies

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226 3.3 Climatic index, hydrological signatures and physiographic characteristics

227 3.3.1 Climatic index

The Aridity Index (\emptyset) is calculated as the ratio of total annual potential evapotranspiration to total annual precipitation. The Aridity Index is related to climatic regimes according to the classification presented by Arora (2002) and Ponce (Ponce et al., 2000), in which the following Aridity Index defines arid, semi-arid, sub-humid, and humid regions ranges: $12 > \emptyset \ge 5$; 5 $> \emptyset \ge 2$; $2 > \emptyset \ge 0.75$; and $0.75 > \emptyset \ge 0.375$, respectively.

The annual frequency of high and low precipitation days is estimated as the number of days in the year with high precipitation (Pr > 5 times the mean daily precipitation) and the number of dry days (Pr < 1 mm/day), respectively. The reported values correspond to the average estimates based on the available data. Similarly, the duration of high precipitation and dry periods per year is measured as the longest consecutive sequence of high precipitation days and dry days. These values also represent the average estimates from the available data.



237 3.3.2 Hydrological signatures

Hydrologic signatures can be key in understanding streamflow patterns by capturing variations like flood peaks and baseflow. They assist in analyzing watershed processes by linking runoff generation and groundwater contributions to physical characteristics. Signatures can improve calibration and evaluation in hydrological modelling by ensuring models replicate key hydrological behaviors. Moreover, they guide regionalization studies, enabling predictions in ungauged basins by transferring knowledge from similar watersheds (McMillan, 2021).

The CAMELS-COL dataset features 11 hydrological signatures for each of the 347 catchments, being computed according to Addor et al. (2018) and described in Table 1, ensuring consistency with previous CAMELS datasets. The annual frequency of high and low flow days is estimated as the number of days in the year with high flow (Q > 9 times the median daily flow) and the number of low flow days (Q < 0.2 times the mean daily flow), respectively. The presented values reflect the average estimates based on the available data. Similarly, the duration of high and low flow events is determined as the longest consecutive number of days with high and low flows. The presented values also represent the average estimates from the available data.

250 3.3.3 Physiographic Characteristics

The estimation of the physiographic characteristics of the catchment is performed using the Whitebox library. The concentration times are estimated using the equations presented by Silveira (2005), with the equivalent slope applied in the estimation. The equivalent slope is used to represent the mean slope of the main river more accurately. On the other hand, the form factor and the compactness index are estimated using the equations presented by Hipólito and Carmo (2017). Table 3 presents the equations used to estimate the physiographic characteristics.



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Geometric characteristic Equation Comments Drainage area (km²) Estimated whit Whitebox library Perimeter (km) Estimated whit Whitebox library $K_c = \frac{0.282 P}{\sqrt{A}}$ Gravelius compactness index $K_f = \frac{A}{L^2}$ Form factor Strength of the main channel (L) Estimated whit Whitebox library **Concentration time** Typically for a small (approximately $T_c = 0.0663L^{0.77}S^{-0.385}$ Kirpich 0.5 km²), forested watershed in a mountainous region (Kirpich, 1940) t is a modification of the Kirpich equation, recommended for watersheds $T_c = 0.160L^{0.64}S^{-0.320}$ with mixed land cover, including some Ven te Chow urban development but predominantly rural, with an area greater than 0.2 km² (Chow, 1964) This method was defined with data $T_c = 0.462 L^{0.50} S^{-0.250}$ from 19 rural basins in the USA, with Johnstone an area range of 64.8-4206 km² (Kaufmann de Almeida et al., 2017). It is commonly used in larger catchments, particularly those with complex flow networks, such as large $T_c = 0.191 L^{0.76} S^{-0.19}$ **Corps Engineers** river basins with multiple tributaries and mixed land cover. (Kaufmann de Almeida et al., 2017) S_i slope of segment i (m/m) $S = \left(\frac{\sum L}{\sum \left(\frac{L_i}{L_i}\right)}\right) \quad ; s_i = \frac{\Delta H_i}{L_i}$ L_i length of of segment i (m) ΔH_i height difference of segment i (m) L Length of main chanel (m)Equivalent slope S Equivalent slope (m/m)

Table 3 Physiographic characteristics estimated

258 259 *L*: Length of the min channel (km), T_c : Concentration time (hours). *A*: Area (km²)

P Perimeter (km)

The compactness coefficient and the form factor are fundamental indicators of a watershed's susceptibility to flooding. The compactness coefficient varies with the shape of the watershed, being higher in more irregular basins (Zävoianu, 1985). Its minimum value is 1, representing a perfectly circular basin, which promotes a greater concentration of surface runoff and,





consequently, a higher risk of flooding. Values between 1.25 and 1.50 indicate a medium to high tendency for flooding, while
coefficients above 1.50 signal a lower susceptibility to these events (Villela and Mattos, 1975). On the other hand, the form
factor expresses the relationship between the basin's shape and the generation of concentrated surface runoff. Values close to
1 indicate more circular basins, where runoff converges quickly to the outlet, increasing the risk of flooding. Lower values
reflect more elongated basins, characterized by less concentrated runoff and a lower tendency for flooding (Villela and Mattos,
1975).

269 The time of concentration equations used in this study are selected because they are the most frequently cited in official technical documents in Colombia (e.g., CORCUENCAS (2014) and Invias (2009)). Due to the variability in the applicability 270 271 of these equations, specific criteria have been established for estimating concentration times. These criteria consider the 272 applicability range of each equation (Table 3) to ensure accurate and context-appropriate calculations. Thus, for areas smaller 273 than 0.2 km², the time of concentration is calculated using the Kirpich equation. For areas between 0.2 and 0.5 km², the time of concentration is determined as the average of the values obtained using the Kirpich and Ven Te Chow equations. For areas 274 between 0.5 and 64.8 km², the Ven Te Chow equation is applied. For areas ranging from 64.8 to 4206.1 km², the concentration 275 276 time is calculated using the Johnstone equation, and for areas larger than 4206.1 km², the concentration time is determined 277 using the Corps of Engineers method. These estimates serve only as an approximation of the time of concentration. Future 278 studies could apply other equations reported in the literature; however, all the results obtained are present in the dataset.

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280 3.3.3.1 Curve Number (CN)

281 The curve number (CN) is widely used in hydrology to estimate direct rainfall runoff, particularly in water resources management studies and stormwater control structure design. Developed by the U.S. Natural Resources Conservation Service 282 (NRCS), the CN is based on soil characteristics, land use, hydrological conditions, and land cover. It plays a crucial role in 283 284 modeling a watershed's hydrological response to precipitation events, helping to assess the amount of water that will infiltrate, store, or run off the surface. The CN was calculated using the ArcHydro Tools "Generate CN Raster" tool (yellow.esri.com -285 (archydro/), designed to generate CN rasters by integrating spatial data on soil, cover, and hydrological parameters. This tool 286 automates the assignment of CN values in each cell, ensuring spatial coherence and homogeneous resolution. The inputs and 287 288 adaptations made for its application are described below.

The first input is the DEM, whose main function is to define the spatial extent and resolution of the output raster. Previously, the DEM underwent a "Fill" correction process to remove anomalies from empty cells, ensuring a continuous and accurate representation of the raster layer. It is important to note that the DEM is not used to adjust CN values as a function of elevation, but only as a spatial reference to align and rasterize the input layers that are in vector format.

The second input corresponded to the soil layer in vector format obtained from IGAC, which was described previously. This dataset describes the drainage capacity and the soil characteristics necessary to classify the different soil types in the



hydrological groups (A, B, C, D). The classification is made using the data available in the attribute table, based on the specifications of the hydrological groups described by the United States Department of Agriculture (Soil Survey Staff, 2016). In this context, hydrological groups are assigned according to the following criteria: well-drained or excessively drained soils correspond to hydrological group A; either or moderately drained to hydrological group B; poorly drained to hydrological group C; while urban areas, bodies of water or soils with very poor drainage are classified in hydrological group D.

The third input corresponded to the land cover in raster format provided by MapBiomas, which is converted to vector format to guarantee compatibility with the "Generate CN Raster" tool. Subsequently, a direct relationship is established between the land cover of MapBiomas and the coverage classes proposed by the United States Army Corps of Engineers in the HEC-HMS manual (Table 4) (USACE, 2023), generated a new land cover denominated in the present work as "Equivalent land cover".

The three previously described variables are processed in the "Generate CN Raster" tool, which provided a raster of 30 meters of spatial resolution, in which each cell represents a CN value adjusted according to the type of land cover and the assigned hydrological group.

	CN by the USACE.	Classifications with the Proposed	nas and USACE	s Between MapBioma	Cover Equivalents	309 Table 4 Land	309
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Land Cover MapBiomas (ID MapBiomas)	USACE land cover (ID USACE)	A	B	С	D
	open_water (11)	98	98	98	98
\mathbf{D} in the second (22)	developed_open (21)	49	69	79	84
River, lake, or ocean (55)	developed_low (22)	57	72	81	86
	developed_med (23)	61	75	83	87
Infrastructure (24)	developed_high (24)	81	88	91	93
Beach,dune and sand spot (23) Other non-vegetated area (25) Not observed (27) Rocky outcrop (29) Mining (30) Glacier (34)	barren_land (31)	78	86	91	93
	deciduous_forest (41)	45	66	77	83
Forest (3)	evergreen_forest (42)	25	55	70	77
	mixed_forest (43)	36	60	73	79
Other non forest formation (13) Wooded sand vegetation (49) Herbaceous sand vegetation (50)	Shrub (52)	55	72	81	86
Grasslands/herbaceous (12)	Grassland (71)	50	69	79	84



Land Cover MapBiomas (ID MapBiomas)	USACE land cover (ID USACE)	A	В	С	D
	pasture (81)	49	69	79	84
Forest plantation (9)					
Mosaic of agriculture and pasture (21)	cultivated_crops (82)	67	78	85	89
Palm oil (35)					
Mangrove (5)	woody watlands (00)	20	50	71	70
Flooded forest (6)	woody_wettailds (90)	30	20	/1	70
Wetland (11)					
Aquaculture (31)	herbaceous_wetlands (95)	30	58	71	78
Hypersaline tidal flat (32)					

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311 3. Results

312 3.1 Catchment and physiographic characteristics

The analyzed catchments exhibit diverse drainage areas. 41% have drainage areas between 28 and 525 km², predominantly located in the mountainous zones of the hydrological regions. 24% of the catchments fall within the 525–1650 km² range, 23,6% have areas between 1650 and 12425 km². The remaining 11.2% correspond to catchments exceeding 12425 km², located predominantly in the Magdalena, Orinoco, and Amazonas regions (Figure 3a).

317 As mentioned, the compactness coefficient and the form factor are used to analyze a watershed's susceptibility to flooding. 318 The results obtained for these indices are consistent and lead to the same conclusion: the studied basins exhibit low 319 susceptibility to flooding. In this context, it is important to highlight that all basins have a compactness coefficient greater than 320 1.5 and a form factor less than 0.5. On the other hand, 16% of the analyzed catchments have concentration times ranging 321 between 4 and 24 hours. Only three catchments present concentration times below 10 hours. However, it is important to note 322 that these are small basins in the Andean region (Station code: 26017040, 35097010, and 35087030). A larger proportion, 323 46%, fall within 24 to 55 hours, while 25% of the catchments exhibit concentration times between 55 and 105 hours, and 13% 324 exceed 105 hours (Figure 3b).

Among the analyzed catchments, 53% exhibit CN values between 60 and 70 (Figure 3c), indicating moderate infiltration capacity and lower runoff generation, typically associated with forested or well-vegetated areas (USDA, 2021). This condition is present in 46% of the stations in the Orinoco region, in 91% of the stations of the Amazon and Caribbean regions, in 53% of the stations of the Pacific region, and in 46% of the stations of the Magdalena region. In catchments where CN values are below 70, forest formations are generally the predominant land cover, which helps explain the observed patterns (Figure 4).





It is also worth highlighting that the Magdalena and Caribbean regions are largely dominated by Inceptisols (Figure 6). When well-preserved—especially under native forest cover—these soils play a key role in regulating streamflow and minimizing runoff. However, their capacity to perform this function is significantly reduced when replaced by agriculture or pasture (Mello et al., 2025). Thus, the prevalence of Inceptisols, along with the conservation of natural vegetation, likely contributes to the lower CN values observed in these regions.

On the other hand, 43% of the catchments have CN values ranging from 70 to 80 (Figure 3e), which are typically associated with agricultural or mixed land cover, where runoff is more pronounced but still not excessive (USDA, 2021). Most of these catchments correspond to the larger basins of the Magdalena region, where land cover is predominantly a mosaic of agriculture and pasture, contributing to higher CN values. Although these findings further reinforce that Colombia's landscape comprises forest formations, agricultural lands, and livestock areas, it also sheds light on the vulnerability of Colombia's headwater to maintain water yield services as agriculture pressure intensifies in the region (Arias and Barragán, 2020).

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Figure 3 Physiographic characteristics of the 347 selected catchments. The size of the circles is proportional to the size of the catchment. A) Range area of the basins, b) Concentration time, c) Curve Number.

345 3.2 Land cover assessment

346 Distinctive land use and cover patterns were identified among the analyzed catchments. Forest formations predominate in 347 43% of the catchments, covering more than 50% of their area. This high forest coverage plays a key role in regulating hydrological processes, favoring infiltration, reducing surface runoff, and promoting water retention, which can contribute to 348 lower susceptibility to flooding (Kan et al., 2019). On the other hand, in 37% of the catchments, agricultural and livestock 349 350 activities occupy more than half of the area principally in the Magdalena region (Figure 4). These land uses tend to increase runoff due to soil compaction that, increases bulk density and penetration resistance and reduces macroporosity, infiltration 351 rates and hydraulic conductivities, amplifying peak flows and influencing sediment transport (Martínez and Zinck, 2004; 352 353 Alegre and Cassel, 1996).



Additionally, 86% of the catchments have less than 20% of their area associated with natural non-forest formations, indicating that these landscapes have a minimal role in the general hydrological dynamics. Non-vegetated areas represent less than 10% of all catchments. The water bodies occupied less of 10% of the basin's area in 99% of the cases. Furthermore, in 93% of the catchments, the land cover was fully identified, while in the remaining 7%, less than 10% of the area could not be classified by the MapBiomas project (see FigureS1 in supplementary material).

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365 3.3 Geology assessment

366 The results show that in 37% of the analyzed basins, sedimentary rocks, formed by weathering pre-existing rocks (Crook and Beck, 2024), constitute more than 50% of the lithological composition (Figure 5). This lithology is present in all 367 hydrological regions and is particularly predominant in the valleys of the Magdalena, Orinoco and Caribbean regions. Due to 368 369 the heterogeneous nature of this group, which includes mudstone, poorly consolidated conglomerates, sandstones with 370 ferruginous and clayey matrices, siltstones, and other sedimentary formations. The second predominant lithology corresponds to plutonic rocks, which comprise more than 50% of the lithological composition in 9% of the analyzed basins (Figure 5). 371 These rocks are mainly found in the valleys of the Magdalena and Orinoco regions. Due to their coarse grain texture and low 372 primary porosity, they have limited infiltration capacity (Watson, 1990). However, when intensely fractured, they can allow 373 374 water infiltration and significantly contribute to aquifer recharge. 375 Metamorphic rocks are the third most common lithology, composing more than 50% of the material in 5% of the analyzed

Metamorphic rocks are the third most common lithology, composing more than 50% of the material in 5% of the analyzed
 basins, and are predominantly found in the valleys of the Magdalena region (Figure 5). These rocks, formed under high pressure
 and temperature, generally have low primary porosity, which limits infiltration (USGS, 2023). In this sense, basins dominated



by metamorphic rocks may have reduced water retention capacity without fracturing, limiting its regularization capacity. On the other hand, 1% of the analyzed stations have a lithological composition dominated by more than 50% volcanic rocks (Figure 5), primarily located in the volcanic region of Colombia, especially near the Puracé, Sotará, Tolima, and Ruiz volcanoes. Characterized by rapid cooling and a fine-grained texture, volcanic rocks typically have low permeability, which may limit infiltration and water storage (Jerram, 2021). These geological characteristics highlight the limited capacity of water storage and release in the volcanic regions (In the Colombian case, primarily in the Magdalena region), which impacts streamflow (mainly in the recession period).

The smallest lithological material in the hydrographic basins includes unconsolidated deposits, hypabyssal rocks, and 385 386 volcaniclastic rocks (Figure 5). Unconsolidated deposits, found across all hydrological regions, are composed of poorly cemented materials, which grant them high porosity and permeability, facilitating infiltration, aquifer recharge, and water 387 388 storage (Wahlstrom, 1974). However, these characteristics depend on the other lithological compositions present. On the other hand, hypabyssal rocks, predominantly located in the Magdalena region, crystallize at intermediate depths and display 389 medium-grained textures. Their permeability can be variable, depending on their degree of fracturing (Macheveki et al., 2020). 390 391 Finally, volcaniclastic rocks found in the volcanic region of Colombia are primarily located in the mountainous areas of the 392 Magdalena region. Composed of volcanic fragments of varying sizes, these rocks may have areas with high initial porosity, 393 though their consolidation over time can reduce infiltration (White and Houghton, 2006).

Although we can have an overview of geology and water storage-release processes from the presented dataset, the geological diversity present in the main hydrographic basins of Colombia underscores the need for individualized analysis of each basin in future studies, intending to account for its specific characteristics.







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400 Figure 5 Geology of the 347 selected catchments. The size of the circles is proportional to the size of the catchment. The lithologies 401 include A) unconsolidated deposits, B) hypabyssal rocks, C) metamorphic rocks, D) plutonic rocks, E) sedimentary rocks, F) 402 volcanoclastic rocks, and G) volcanic rocks.

403



405 3.4 Soil Assessment

The soil composition analysis in Colombia's watersheds revealed that the predominant soil types are Andisols, Entisols, and Inceptisols (Figure 6). The other soil types generally cover less than 10% of the watershed area (e.g., Alfisols and urban areas). They are sometimes absent, as with Spodosols and soils associated with open-pit coal mines. The spatial distribution of these soil types, distinct from Andisols, Entisols, and Inceptisols, is presented in the item "Spatial distribution of the Secondary soil typologies" of supplementary material (Figure S2).



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Figure 6 Main Soil Typologies of the 347 Selected Catchments: The Circle Size is Proportional to the Catchment Area: A) Inceptisols, B) Entisols, C) Andisols.

Andisols, characterized by their volcanic origin, exhibit good permeability and drainage capacity due to their fine and porous texture derived from volcanic ash (Uehara, 2005). It was identified that 22% of the watersheds have more than 40% of their area covered by Andisols, predominantly in areas near the volcanic region of the Andes, where the dominant lithology consists of volcanic and volcaniclastic rocks. Additionally, 34% of the watersheds have between 10% and 40% of their area occupied by Andisols, mainly in the Pacific and Magdalena hydrological regions.

Entisols, soils with limited pedological development that can exhibit variability in drainage capacity depending on the parent material, are commonly found in alluvial plains and areas with high erosion (Eswaran and Reich, 2005). It was observed that 23% of the analyzed watersheds have between 10% and 40% of their surface covered by this soil type, with a broad spatial distribution across all hydrological regions.

Finally, Inceptisols, which represent an intermediate stage in soil development, and whose drainage capacity depends on the underlying lithology (Foss et al., 1983), are the most prevalent in Colombia. It was observed that 39% of the studied watersheds have more than 40% of their surface covered by Inceptisols, with a concentrated distribution in the Magdalena, Pacific, and Caribbean regions, as well as in the western sector of the Amazon and Orinoco regions, near the Andes Mountain range.





429 Entisols and Inceptsols are poorly developed soils, with shallow horizons and steep slopes (Idaho, 2025). Therefore, they 430 are most present in the headwater region of the Andes Mountains. These soils exhibit specific hydropedological characteristics 431 that influence water infiltration, retention, and release. For instance, Inceptsols contribute significantly to water yield, but only 432 when well-preserved (Mello et al., 2025) because root systems and organic matter accumulation enhance soil hydraulic 433 properties, improving infiltration and percolation. In contrast, Entisols are shallow soils close to the parent material, which limits its water storage capacity (except in basins dominated by snowmelt processes). This makes them more susceptible to 434 435 runoff formation and erosion (Kumar et al., 2000). On the other hand, Andisols improve water storage and release (Uehara, 436 2005), supplying streamflow in recession periods and increasing overall water yield. In summary, the hydrological responses 437 of basins with predominance of Entisols and Inceptisols are more sensitive to land use and cover changes.

438 3.5 Climatological Assessment

The precipitation patterns of Colombian catchments exhibit significant spatial variability. Notably, river basins located in the western Caribbean and Pacific regions register the highest daily precipitation values (Pr > 10 mm/day). In the Amazonian catchments, the mean daily precipitation ranges between 8 mm/day and 10 mm/day. In comparison, most basins in the Magdalena region show daily precipitation values varying between 4 mm/day and 6 mm/day (Figure 7a).

443 Potential evapotranspiration is directly linked to temperature. In this regard, the catchments in the Pacific, Caribbean, 444 Orinoco and Amazon regions, as well as the large basins of the Magdalena region, present the highest potential 445 evapotranspiration values, ranging from 9 to 13 mm/day (Figure 7b). Temperature, on the other hand, shows an inverse 446 relationship with altitude. The lowest mean daily temperatures (8°C - 17°C) are recorded in the high-mountain catchments of 447 the Magdalena, while catchments with outlet points in the inter-Magdalena valleys predominantly experience temperatures 448 between 17°C and 23°C. In contrast, the catchments in the Pacific, Caribbean and Amazon regions, as well as the large basins 449 of the Magdalena region, register temperatures exceeding 23°CFinally, the highest temperatures are observed in the Orinoco region, where values exceed 29 °C (Figure 7c). 450

The analysis of the aridity index reveals that 74.7% of the analyzed catchments exhibit a subhumid climate regime, encompassing catchments from the Magdalena, Orinoco, and Amazon regions. Meanwhile, the eastern flank of the Andes and the Caribbean region display a semiarid climate regime, accounting for 20.7% of the analyzed catchments. Lastly, only 4.6% of the catchments fall within a humid climate regime, predominantly located on the western flank of the Caribbean and Pacific regions (Figure 7d).

Catchments with the highest occurrence of extreme precipitation days (\geq 5 times the mean daily precipitation) are concentrated in the mountainous areas of Magdalena region thus as Orinoco region. In the case of Magdalena, this pattern may be associated with orographic effects, as mountain ranges such as the Eastern Cordillera act as barriers to the southeast trade winds, promoting the formation of orographic rainfall (Ruiz et al., 2017; Espinoza et al., 2015). In contrast, the northern part of the country exhibits fewer extreme precipitation days, possibly due to the reduced influence of the Andes' orographic effects. Additionally, a direct spatial correlation is observed between the number of high-precipitation days and their duration.



462 Catchments with a high number of high-precipitation days also tend to experience longer consecutive sequences of such events463 (Figure 7 e and f).

The analyses also reveal that the Andes Mountain range catchments experience the highest number of dry days per year (>190 days/year). This can be attributed to the orographic shadow effect, where moist winds from the Atlantic and Pacific lose much of their humidity as they ascend the mountain slopes, reducing precipitation in the high-altitude areas of the Colombian Massif. Another relevant factor is the oscillation of the Intertropical Convergence Zone (ITCZ), which shifts north or south during certain periods of the year, reducing the influence of rain-generating systems in the Magdalena region (Loaiza Cerón et al., 2020).

Catchments with outlets in non-mountainous areas experience fewer dry days, ranging from 100 to 190 days. Notably, the country's largest river basins record the lowest dry days (47–100 days/year) (Figure 7g). River basins in the Pacific region experience the fewest consecutive dry days, likely due to the region's high and more evenly distributed rainfall throughout the year. In the Magdalena, Orinoco, and Amazon regions, the number of consecutive dry days ranges from 9 to 12. However, basins located in the mountainous areas of the Andes experience a higher number of consecutive dry days, exceeding 18 days (Figure 7h).







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Figure 7 Climatic index of the 347selected catchments. The size of the circles is proportional to the size of the catchment. A) Mean
 daily precipitation in the catchment, b) Mean daily ETo in the catchment, c) Mean daily temperature in the catchment, d)
 Catchment aridity index, e) Mean frequency of high precipitations days, f) Mean duration of high precipitations days, g) Mean
 frequency of low precipitations days, h) Mean duration of low precipitations days.



485 3.6 Hydrological Signatures

Figure 8a shows that 58% of the basins have an average daily flow of less than 3.3 mm/day, predominantly located in the Magdalena region. In contrast, basins in the Orinoco and Amazon regions exhibit average daily flows ranging between 3.3 and 6.3 mm/day, while those in the Pacific region present the highest values (> 6.3 mm/day). These findings indicate a relationship with daily precipitation, meaning regions with higher precipitation also tend to have higher streamflows. On the other hand, Q5 and Q95 follow a similar spatial pattern to the mean daily flow, with higher values in the Pacific basin and lower values in the Magdalena basin (Figure 8b and Figure 8c).

492 The runoff coefficient predominantly ranges between 0.3 and 0.7, with 56% of the analyzed basins falling within this 493 range. In the Magdalena region, 57% of the catchments exhibit runoff coefficients between 0.4 and 0.7. The predominant runoff coefficient ranges from 0.5 to 0.7 in the Caribbean region, with 33% of the basins reporting values within this range 494 (Figure 8d). In the Orinoco region, 35% of the basins recorded runoff coefficients greater than 1, while this proportion increases 495 496 to 38% in the Pacific region and 41% in the Amazon region. In the Orinoco and Amazon regions, these basins are primarily located on the western slopes of the Andes Mountain range. Additionally, 37% of the basins in the Orinoco region present 497 runoff coefficients between 0.5 and 0.7, while 32% of the Amazon basins report runoff coefficients between 0.7 and 1.0. 498 499 Runoff coefficients exceeding 1 may be linked to an underestimation of precipitation by the CHIRPS satellite, as reported by 500 Cavalcante et al. (2020) and Paredes-Trejo et al. (2017).

501 One third of Colombia's river basins (34%), mainly in the Amazon and Orinoco regions, exhibited a streamflow elasticity 502 ranging from 1.5 to 2.0, indicating that a 1% variation in precipitation could result in a 1.5% to 2.0% change in streamflow 503 (Anderson et al., 2024). In the Orinoco region, elasticity ranged from 1.0 to 2.5. In the Magdalena region, values varied between 504 0.5 and 4.0, with a predominance of elasticities between 1.0 and 2.5 observed in 71% of the region stations. Meanwhile, in the 505 Caribbean region, elasticities fluctuated between 0.5 and 4.0, with a predominant range of 1.0 to 2.5. In this region, 15% of 506 the stations recorded elasticities between 1.0 and 1.5, 37% between 1.5 and 2.0, and 30% reported values above 2.0 (Figure 507 8e).

Regarding the flow duration curve, 53% of the basins exhibit a slope between 0.6 and 0.8, predominantly in the Magdalena, Pacific, and Caribbean regions. This suggests a balance between base flow and response to precipitation events, considering that a high slope (>0.8) indicates a flashy flow regime, while a low slope suggests a more damped flow (Chouaib et al., 2018; Sawicz et al., 2011). In contrast, basins with high slopes (>0.8) are found in the Amazon and Orinoco regions, as well as in the eastern flank of the Magdalena region, within the Andean Mountain range massif (Figure 8f).

513 In Colombia, 76% of river basins have a Base Flow Index (BFI) between 0.6 and 0.8, suggesting that groundwater 514 resources play a crucial role in the stability, availability, and regulation of surface water sources. Furthermore, a correlation is 515 observed between land cover and BFI, as basins with an index above 0.6 are predominantly covered by forests, which occupy 516 more than 30% of the total basin area (Figure 8g).



517 Considering that most of the country's basins exhibit an elasticity index greater than 1 and runoff coefficients 518 predominantly below 0.7, any alteration in the precipitation regime, deforestation, or overexploitation of aquifers is expected 519 to compromise Colombia's water stability and availability. Furthermore, the slope of the flow duration curve reinforces this 520 vulnerability, especially in the Orinoco and Amazon basins, where approximately 52% of the country's groundwater reserves 521 are concentrated (Aranguren-Díaz et al., 2024) and where high slopes of the flow duration curve (>0.8) prevail. This suggests 522 a more variable flow regime, leading to greater susceptibility to changes in water availability in response to any alteration in 523 precipitation, land use change, or aquifer overexploitation.

524 On the other hand, 61% of the basins analyzed did not record any high-flow days, while another 14% experienced only 525 one high-flow day per year. About 15% of the basins had 2, 3, or 4 high-flow days per year, while the remaining 15% reported 526 between 4 and 30 days of high-flow. It is observed that basins with more than 2 high-flow days per year are predominantly 527 located on the eastern flank of the Andes Mountain range, in the transition zone between the Magdalena and Orinoco regions. 528 A correlation is evident regarding the duration of these events with the total number of high-flow days. In this regard, 14% of 529 the basins recorded only one consecutive day of high flows, two consecutive high flows account 14% of the basins as well, 530 also 14% experienced between 3 and 5 consecutive days, and 21% of the basins registered events lasting more than 5 531 consecutive days (Figure S3 of supplementary material).

The number of low-flow days is higher than that of high-flow days. In this regard, basins in the western part of the country predominantly exhibit fewer low-flow days (0–5 days per year) compared to those in the eastern region, where values exceed 5 days per year. The highest number of low-flow days is primarily observed in the upper reaches of the Andes, particularly along the mountain range near the hydrological divide between the Magdalena, Orinoco, and Amazon regions (Eastern Cordillera), as well as in the Amazon and Orinoco regions. Additionally, a relationship is observed between the maximum number of consecutive low-flow days and the number of low-flow days per year, suggesting that basins experiencing frequent low-flow conditions also tend to have prolonged periods of consecutive low flows (Figure 8j and Figure 8k).

539 The observed hydrological signatures shed light on the vulnerability of most Colombian basins to changes in precipitation. 540 This aligns with the previous statement regarding soil properties and how basin responses are driven by the land use and cover 541 conditions. High elasticity values (mostly between 1.5 and 2.5) represent the sensitivity of streamflow to precipitation 542 variability (Anderson et al., 2024). This is reinforced by the BFI values ranging from 0.6 to 0.8, which are lower than those 543 observed in basins overlying productive aquifers (Collischonn and Dornelles, 2013). Thus, water storage contributions to 544 streamflow appear to be constrained by soil water-holding capacity, depth, and release processes. This limited storage capacity 545 is reflected in the slope of the flow duration curve (ranging mainly from 0.4 to 1.0), highlighting the basins' difficulty in 546 sustaining streamflow (lower regulation). In summary, Colombian basins seem more sensitive to climate variability than to 547 hydrological processes (e.g., water storage and release), i.e., climate change is likely to intensify extreme events such as floods 548 and droughts in these regions.







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Figure 8 Hydrological signatures of the 347 selected catchments. The size of the circles is proportional to the size of the catchment.
 a) Mean daily discharge, b) Q5 (5th percentile daily flow), c) Q95 (95th percentile daily flow), d) Runoff ratio, e) Streamflow
 precipitation elasticity, f) Slope of the FDC, g) Baseflow index, h) Frequency of high-flow days, i) Duration of high-flow days, j)
 Frequency of low-flow days, k) Duration of low-flow days.



556 4. Water Resources and Climate Change in Colombia

557 Colombia's agriculture and energy production are particularly sensitive to hydrological dynamics, making them 558 dependent on both climate variability and water availability. For instance, coffee and rice crops, which are vital to both food 559 security and export revenues, rely on rainfall seasonality and irrigation schedules. As climate change increases the rainfall 560 spatiotemporal variability in Colombia, water availability in both time and space is likely to shift, impacting harvesting schedules and crop yields. Moreover, approximately 70% of Colombia's power generation comes from hydroelectricity (IEA, 561 562 2023), making it vulnerable to streamflow variability. In the Magdalena River Basin, projections under RCP4.5 and RCP8.5 563 scenarios show reduced dry season discharge in the headwater regions and increased inundation downstream due to intensified 564 rainfall (Munar et al., 2023). This pattern is confirmed in the IDEAM (2024), which assessed the impacts of climate change across Colombia based on the CMIP6 and different emission scenarios. On the other hand, drying trends are expected in 565 February (-0.60 mm/year) in the Caribbean region, while some departments are likely to experience increases up to 1.61 566 567 mm/year in November (Arregocés et al., 2024). However, the annual precipitation is likely to reduce by up to 44% under highemission scenarios (IDEAM et al., 2024). Hydrological signatures (such as BFI and elasticity) highlight the sensitivity of these 568 watersheds to climate change, as they are highly sensitive to precipitation, which in turn impacts streamflow (McMillan, 2020). 569 570 BFI values indicate low to medium capacity of streamflow regularization (Figure 9g), largely due to the predominant presence 571 of Inceptisols and Entisols, as well as the presence of porous and fractured aquifers (Figures 6 and 7). Therefore, the Magdalena 572 and Caribbean regions are highly sensitive to climate change as their hydrological responses are more influenced by weather 573 variability than hydrological stability.

574 Water availability is also dependent on the Páramo ecosystems, such as those in the Chingaza National Park. The 575 Páramo ecosystem is over 3,000 m a.s.l and mostly in areas of Andisols (Figure 6). Andisols have good permeability and 576 elevated porosity (Uehara, 2005), which together with the Páramo improves organic matter accumulation (due to low temperatures in high altitudes), increases water infiltration and storage, and regulates baseflow, configuring it as "natural 577 578 sponges" (Herzog et al., 2011). Moreover, this ecosystem can intercept cloud and mist, increasing water input to the soil as 579 occult precipitation (Cárdenas et al., 2017). However, this ecosystem is highly sensitive to warming and dry conditions (Cresso 580 et al., 2020). Climate projections suggest that up to 52% of the Páramo area could become unsuitable, undermining ecosystem 581 services such as groundwater recharge and carbon sequestration (Cresso et al., 2020). As much of the watersheds in the Andean 582 region have elasticity > 1.0 (Figure 9e), they are sensitive to precipitation, and a slight change in rainfall can lead to intensified 583 droughts or floods, as Páramos are likely to lose their ability to produce water (Cresso et al., 2020).

Mean temperatures are projected to increase by 0.7 to 2.5 °C under SSP2-4.5 in the mid-century across Colombia (IDEAM et al., 2024). The official Colombia's report (IDEAM et al., 2024) also projects an increase up to 5.0°C in the five hydrological regions of Colombia (Pacific, Magdalena, Orinoco, Amazon, and Caribbean) by the end of the century. Although this can be less important than precipitation changes, due to the high dependence of Colombia's watershed on precipitation (indicated by hydrological signatures), this increment potentially increases atmospheric water demand and pressures water



availability. Irrigation schedule and demand changes with greater evapotranspiration to keep crop yield as observed in a densely irrigated area in Brazil (Rodrigues et al., 2024). In this sense, the projected precipitation variability and air temperature increment will likely increase water conflicts across Colombia, mainly in the watersheds that are dependent on the Andean and Páramo ecosystems' services. Moreover, the country's hydroclimate is heavily influenced by the El Niño–Southern Oscillation (ENSO), which may cause recurring anomalies such as intense droughts during El Niño and excessive flooding during La Niña (Rodriguez-Espinosa et al., 2025). However, the effects of large-scale climate anomalies in Colombia are still unclear, and more studies are necessary to improve our understanding of their impacts and the areas most affected.

596 CAMELS-COL provides enough information to improve continental-scale understanding on water resources response 597 to climate change, decreasing uncertainties of current models, especially in underserved regions such as the Amazon and 598 Orinoco basins, where data is sparse. Moreover, this dataset sheds light on the sensibility of Colombia's watersheds to weather 599 variability, which can face more frequent floods and droughts depending on the increase or decrease of precipitation. This 600 information can also support a coordinated national strategy for adaptive water governance, backed by scientific monitoring 601 and updated design standards, is essential to reduce future climate risks. Moreover, it adds to the CAMELS-CH (Höge et al., 602 2023) and CAMELS-BR (Chagas et al., 2020) and enables advances on the South America hydrological understanding and 603 response to climate change.

604 5. Conclusions

505 So far, large-sample hydrological studies in Colombia have lacked a comprehensive and easily accessible dataset. This 506 study introduces CAMELS-COL, a new dataset that provides streamflow time series for 347 catchments nationwide. In 507 addition to streamflow data, CAMELS-COL includes climatic series, physiographic characteristics, climatic indices, 508 hydrological signatures, land cover, geology, and soil properties, primarily derived from official governmental sources.

The analysis of Colombian catchments highlights the diverse physiographic, hydrological, geological, and climatological factors influencing water dynamics across the country. In this sense, the analyzed basins exhibit significant variability in drainage area, with 41% ranging between 28 and 525 km², predominantly located in the mountainous zones of the Magdalena, Caribbean, Amazon, and Orinoco regions. The compactness coefficient and form factor suggest that most evaluated Colombian basins have low susceptibility to flooding. Additionally, basin concentration times vary widely, from less than 10 hours to over 105 hours.

Forest formations predominate in 43% of the catchments, and Agricultural and livestock land cover dominate 37%. In contrast, non-vegetated areas and water bodies have a less prominent role in the hydrological processes of the analyzed basins, as they represent less than 10% of the land cover total basin area in most cases. Regarding geological characteristics, sedimentary rocks dominate the lithological composition in 37% of the catchments. Plutonic rocks are the second most prevalent lithology, constituting more than 50% of the lithological composition in 9% of the catchments. Finally, metamorphic and volcanic rocks account for over 50% of the basin area in 5% and 1% of the analyzed catchments, respectively.



Precipitation patterns in Colombian river basins exhibit high spatial variability. The highest daily values are observed in the Caribbean and Pacific regions, while the Magdalena region experiences lower daily precipitation. Potential evapotranspiration follows a direct relationship with temperature, with the highest values recorded in the Caribbean, Pacific, Amazon, and larger basins of the Magdalena.

An analysis of the aridity index reveals that 74.6% of the basins have a subhumid climate regime, while 20.7% exhibit semiarid conditions and only 4.6% are classified as humid. The distribution of extreme precipitation days indicates that the Magdalena basins, particularly those in mountainous areas, experience the highest frequency of heavy rainfall events while recording the greatest number of dry days.

From a hydrological perspective, the lowest mean daily streamflows are found in the Magdalena basins. In contrast, the highest values occur in the Pacific, Amazon, and Orinoco regions, aligning with precipitation distribution patterns. The runoff coefficient generally ranges from 0.3 to 0.7 across most basins; however, values exceeding 1 are observed in the Pacific, Orinoco, and Amazon regions, possibly due to underestimations in satellite-based precipitation data.

Flow elasticity in response to precipitation is highest in the Amazon and Orinoco, indicating a greater streamflow sensitivity to rainfall changes. The slope of the flow duration curve suggests that basins in the Orinoco and Amazon regions exhibit greater hydrological variability, making them more vulnerable to shifts in water availability caused by changes in precipitation patterns, deforestation, or overexploitation of aquifers. The Base Flow Index highlights the crucial role of groundwater in stabilizing and regulating surface water, particularly in basins with extensive forest cover. Finally, high-flow days are most frequent in the transitional zone between the Magdalena, Orinoco, and Amazon regions. In contrast, low-flow days are more predominant in the eastern part of the country.

The CN values of the analyzed catchments range from 50 to 82, with 54% falling between 60 and 70, indicating moderate infiltration capacity and relatively lower runoff generation. These catchments are mainly found in the Amazon, Orinoco, Pacific, and Caribbean hydrological regions, as well as in the mountainous areas of the Magdalena region, where the predominance of forest formations as the primary land cover explains the observed CN values. Conversely, another 43% of the catchments exhibit CN values between 70 and 80, typically associated with agricultural or mixed land cover, where runoff is more pronounced but not excessive. Most of these catchments correspond to the larger basins of the Magdalena region, where land cover consists mainly of a mosaic of agriculture and pasture, contributing to higher CN values.

The CAMELS-COL dataset represents a significant advancement in large-sample hydrology for Colombia, providing a standardized and accessible resource for analyzing the country's diverse hydrological, climatic, and physiographic characteristics. It enhances the benchmarking, calibration, and validation of hydrological models while supporting flood and drought monitoring efforts. Additionally, it contributes to climate change adaptation planning at the departmental level and enables global and regional hydroclimatic comparisons. Future research should integrate this dataset with dynamic hydrological models to improve water resource management strategies and assess the long-term impacts of climate change on Colombian river basins.





654 6. Data availability

The CAMELS-COL dataset is freely available at <u>https://doi.org/10.5281/zenodo.15554735</u> (Jimenez et al., 2025). The files provided are (i) the Geologic_characteristics, Land cover characteristics, Soil characteristics, Climatic indices, Hydrological signatures and Physiograpic Characteristics in *.xlsx files (ii) the daily time series in zip files, (iii) the catchment boundaries of the catchments in zip file (iv) a readme file.

659 7. Author contributions

DAJ, BMB, and AFR initiated the investigation. JEM and AAD estimated the catchments' Curve Numbers (CN), while
PHBS calculated the hydrological signatures. DAJ, AAD, and BQ performed the analysis of climatological information. FR
and BMB designed the conceptual framework. All authors contributed to the interpretation of the results, the discussion of the
findings, and the editing of the manuscript.

664 8. Competing interests

- 665 The authors declare that they have no conflict of interest.
- 666

667 9. Acknowledgments

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673 **10. References**

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