



Long-term	hourly stream-water flux data to study the effects of forest management on
solute tran	isport processes at the catchment scale
Heye R.	Bogena ^{*1} , Frank Herrmann ¹ , Andreas Lücke ¹ , Thomas Pütz ¹ and Harry
Vereecken	1
Authors' a	addresses:
1: Forschu	ingszentrum Jülich GmbH, Institute of Bio- and Geosciences, Agrosphere
Institute (II	BG-3), Wilhelm-Johnen-Straße, 52425 Jülich, Germany
*Correspo	nding author: Forschungszentrum Jülich GmbH, Institute of Bio- and
Geoscience	es, Agrosphere Institute (IBG-3), Wilhelm-Johnen-Straße, 52425 Jülich,
Germany, l	h.bogena@fz-juelich.de





21 Abstract

A substantial body of knowledge concerning the functioning of catchments has been derived 22 23 from the quantification of solute and suspended matter fluxes in rivers. The Wüstebach catchment is a hydrological observatory that is part of the German TERENO (Terrestrial 24 25 Environmental Observatories) network. In 2013, the Eifel National Park undertook a partial deforestation of the spruce forest with the objective of facilitating the regrowth of a natural 26 27 forest. This data paper presents 16 years of estimated hourly stream-water flux data of nine continuously monitored macro- and micronutrients, as well as dissolved ionic aluminum 28 29 and dissolved organic carbon (DOC), along with the measured solute concentrations and discharge rates observed in the Wüstebach catchment (from 2010 to 2024). 30

To estimate hourly stream-water fluxes from weekly manual grab samples and event 31 autosampler data, we employed the R software package LOADFLEX, which implements a 32 33 number of solute prediction methods, including regressions, interpolations, the period-34 weighted approach, and the more recently developed composite method. A comparison of 35 the predicted nitrate concentrations with hourly nitrate reference data was conducted to assess the optimal prediction approach for the Wüstebach catchment. The analysis showed 36 37 that the composite model is best suited to calculate the nitrate fluxes. Accordingly, this model was selected to calculate the fluxes of all considered macro- and micronutrients, 38 39 dissolved aluminum and DOC. Flux data were compiled in the same way for a neighboring reference catchment with similar characteristics but without clear-cutting, in order to 40 identify the effects of deforestation and afforestation on the cycling and transport of 41 nutrients. We anticipate that this comprehensive data set will facilitate new insights into the 42 influence of deforestation and afforestation on solute fluxes at the catchment scale. The 43 dataset, entitled "Wüstebach data paper: Long-term hourly solute flux data 2010-2024", is 44 shared Forschungszentrum Jülich: https://doi.org/10.26165/JUELICH-45 via DATA/AKAMNQ (Bogena and Herrmann, 2025). 46

- 47
- 48
- 49

50 Keywords

51 Wüstebach catchment; deforestation; afforestation; stream-water flux data; solute
52 prediction models; composite model; R software





53 1 Introduction

Stream discharge and dissolved concentrations of anions, cations and DOC determine how 54 55 much solute mass is exported from catchments (Wang et al., 2024). In recent decades, a substantial body of research has been conducted to examine the impact of deforestation and 56 57 extensive forest die-off on solute transport at the catchment scale, e.g. Keller, 1970; Feller, 2005; Mikkelson et al., 2013; Georgiev et al., 2021; Vilhar et al., 2022; Rajwa-Kuligiewicz 58 59 and Bojarczuk, 2024. There is evidence that forest dieback is increasing due to climate change (Deuffic et al., 2020). For instance, following the successive 2018-2019 hot 60 61 droughts, an extensive mortality event was observed in forests in Germany (Obladen et al., 2021; Xu et al., 2025), with serious consequences for the water quality of rivers and 62 reservoirs (Kong et al., 2022). Nevertheless, there is still a notable lack of comprehensive 63 long-term data sets of stream-water fluxes with high temporal resolution covering all 64 important anions and cations and accompanying studies of the effects of forest dieback and 65 clear-cut on the water and nutrient balance in a holistic approach (Mollenhauer et al., 2018). 66

To address this research gap, a deforestation experiment was conducted in 2013 in the Wüstebach catchment (Bogena et al., 2015) as part of the German TERENO (Terrestrial Environmental Observatories) initiative (Zacharias et al., 2011, 2024) to investigate the solute flux response to the clear-cutting operation. Long-term data before and after forest clear-cut, in combination with water quality models, improve research into the effects of forest dieback and the subsequent recovery of forests (Chen et al., 2024; Musolff et al., 2024).

Solute flux estimates are an indispensable element of whole-catchment manipulation experiments for many decades (e.g. Likens et al., 1970). The total solute flux for given period, Φ_T , is the convolution of solute concentration C and discharge Q over time t (Aulenbach and Hooper, 2006):

78

$$\Phi_T = \int \mathcal{C}(t)Q(t) \, dt \tag{1}$$

Solving (1) requires continuous measurements of solute concentrations and discharge. While continuous discharge data (e.g. at 10-min resolution) is often available, solute concentrations are typically measured on a weekly or event basis as was the case in this study. A plethora of methods have been proposed to determine Φ_T from discontinuous data including straightforward period-weighted means and linear and nonlinear statistical models (Birgand et al., 2010; Cox et al., 2008; Preston et al., 1989; Worrall et al., 2013). These methods are employed for the purpose of interpolating the actually measured solute





86 concentrations to a higher temporal resolution (e.g. hourly). More sophisticated flux calculation methodologies have been incorporated into software applications, including 87 88 FLUX (Walker, 1996), LOADEST (Runkel et al., 2004), and EGRET (Hirsch and De Cicco, 2015). The most comprehensive solute flux calculation software package to date is 89 LOADFLEX (Appling et al., 2015), which includes several estimation models and an 90 algorithm for estimating the prediction uncertainty. LOADFLEX has been used already for 91 92 solute flux calculations in several studies, e.g. Brunet et al., 2021; Coble et al., 2018; Harley et al., 2023; McDowell et al., 2019. 93

94 In this data paper, we provide 16 years of hourly data on the fluxes of nine regularly monitored macro- and micronutrients, as well as dissolved ionic aluminum and DOC, in the 95 Wüstebach catchment. This work also updates and expands on earlier solute flux 96 97 calculations in the Wüstebach that were based only on simple weekly averages (Płaczkowska et al., 2022). To determine the most appropriate method for determining 98 continuous solute concentrations, four well-known methods (i.e. rectangular interpolation, 99 100 two regression models and a composite model) were applied to the example of nitrate and 101 compared with high-resolution nitrate measurements taken in the Wüstebach stream over a period of two years. We also compared the results of the DOC flux predictions using the 102 four methods. Finally, the composite model was used to determine all solute fluxes for the 103 Wüstebach catchment and a neighboring reference catchment that was not affected by the 104 clear-cutting. 105

The presented high-resolution flux data are of great importance as they reflect the effects of clear-cutting measures as well as climate change on the biogeochemical cycles in river catchments. Furthermore, these long-term data can also be interpreted in the context of the environmental observation data available within TERENO (e.g. land-surface-atmosphere exchange fluxes, soil water storage and chemistry, as well as vegetation properties).

111

112 **2** Site description

The experimental Wüstebach headwater catchment (39 ha; WU14, Figure 1) belongs to the Lower Rhine/Eifel Observatory of the TERENO network (Bogena et al., 2018; Zacharias et al., 2024) and is located in the Eifel National Park (50°30'16"N, 06°20'00"E). The area is predominantly covered by Norway spruce (Picea abis L.), with altitudes ranging from 595 m to 628 m above sea level and an average slope of 3.6 percent. The region belongs to the temperate oceanic climate (Cfb) according to the Köppen climate classification system,





119	with an average annual temperature of approx. 7 $^{\circ}\mathrm{C}$ and an average annual precipitation and
120	runoff of approx. 1200 mm and 700 mm, respectively (Graf et al., 2014). The subsoil is
121	composed of Devonian slate, which is overlayed by a periglacial solifluction layer with a
122	maximum depth of 2 m. Cambisols and Planosols are typically found at groundwater distant
123	sites, whereas Gleysols and Histosols have developed as a result of groundwater influence
124	in the riparian zone of the valley (Figure 1).





126

Figure 1: Map showing the main soil types of the Wüstebach and reference catchments and the deforestation area. The contour lines and the locations of the climate station WU_K_2 and the discharge stations WU10, WU14 and WU17 are also shown (geographical coordinates of the stations are given in Table A2).

131

A smaller tributary catchment (11 ha) with similar characteristics serves as an undisturbed reference catchment for the clear-cut experiment (WU17, Figure 1). The similarity, which is important to investigate the impact of deforestation on the hydrological system, was confirmed by comparing the runoff dynamics of both catchments before deforestation over a period of more than three years (Wiekenkamp et al., 2016a).

Over the past two decades, the Wüstebach catchment has been the subject of intensiveresearch, with a particular focus on hydrological (Bogena et al., 2010; Rosenbaum et al.,





2012; Graf et al., 2014; Stockinger et al., 2014; Wiekenkamp et al, 2016b) and
biogeochemical processes (Bol et al., 2015; Gottselig et al., 2017; Weigand et al., 2017).
From August to September 2013, 8.6 hectares of the Wüstebach catchment area were
cleared in order to allow natural succession to a mixed forest to occur (Figure 1).

143 More recently, studies investigated also the effects of the deforestation on hydrological and 144 land surface processes in the Wüstebach catchment (Ney et al., 2019; Wiekenkamp et al., 2016a, 2019; Bogena et al., 2021; Wang et al., 2021; Robinson et al., 2022; Wang et al., 145 2022; Heistermann et al., 2022; Płaczkowska et al., 2022, 2024). The clear-cut measure has 146 147 had a significant impact on the hydrological system of the Wüstebach catchment, affecting several key hydrological processes (Wiekenkamp et al., 2019). Following the clear-cutting, 148 evapotranspiration was initially reduced by approximately 50% and this decline was 149 observed to subside and return towards pre-cut values during the first two years (Ney et al., 150 2019). Wiekenkamp et al. (2016a) observed that the reduction in evapotranspiration resulted 151 in an increase in soil water content within the deforested catchment area during the initial 152 two-year period following deforestation, particularly during the summer months. 153 Furthermore, Wiekenkamp et al. (2019) revealed that the partial clear-cut resulted in an 154 increase in the occurrence of preferential flow within the deforested area. In consequence, 155 this has led to an increased frequency of high flows. Wang et al. (2022) and Robinson et al. 156 (2022) investigated the deforestation effects on nutrient transport and found that following 157 the clear-cut operation, there was a notable increase in DOC and nitrate concentration levels 158 in stream water. For instance, for DOC, a 60% increase was observed in the Wüstebach 159 stream within the first two years after deforestation (Wang et al., 2022). In addition, 160 Robinson et al. (2022) showed that stream water nutrient concentration levels increased 161 162 following the clear-cutting operation, reaching a peak approximately two to three years later and subsequently returning to pre-deforestation levels after approximately five years. 163

164

165 **3 Data and methods**

166 *3.1 Precipitation and discharge measurements*

Precipitation is measured at the climate station WU_K_2 located in the clear-cut area (Figure 1) using a weighing rain gauge (Pluvio², Ott Hydromet GmbH, Kempten, Germany). Precipitation data which were not available at WU_K_2 (e.g. before the clearcut) but used in this work for gap-filling came from the official climate station Kalterherberg (German Weather Service, DWD), which is located 8 km to the west at 535 m a.s.l. In a





previous study, it was demonstrated that the precipitation data from climate stationKalterherberg are representative for the Wüstebach catchment (Graf et al., 2014).

Continuous stream discharge measurements were conducted at the outlets of the Wüstebach 174 and control streams (WU14 and WU17) at a temporal resolution of 10 min. Data from 175 176 another Wüstebach discharge station (WU10) was used for gap-filling. The low flow measurements are taken using a V-notch weir, while the medium to high flow measurements 177 178 are taken using a Parshall flume (ecoTech Umwelt-Messsysteme GmbH, Germany). Both weir types employ a pressure sensor (PDL, ecoTech Umwelt-Messsysteme GmbH, 179 180 Germany) to measure water level fluctuations at ten-minute intervals. The precision of these measurements is $\pm 0.25\%$. The water levels from the two discharge gauging weirs were 181 converted into discharge values using well-established stage-discharge relationships. 182 Subsequently, the two discharge datasets were combined based on the water levels 183 measured in the V-notch weir: If the water level is below 5 cm, the discharge values derived 184 from the V-notch weir are utilized. Conversely, when the water level is above 10 cm, the 185 discharge values from the Parshall flume are employed. The weighted mean of both 186 discharge values was then calculated for water levels between 5 and 10 cm (Bogena et al., 187 2015). 188

The measured discharge time-series have several gaps for technical reasons. Since a 189 complete discharge record is necessary for solute load calculation, the data gaps were filled 190 using a multi-tiered approach. First, the data gaps at WU14 were filled with the help of the 191 192 data from WU10. The ratios of the discharge values at the two discharge stations were calculated for all times for which pairs of measurements were available. These ratios vary 193 within a characteristic range. For periods without available discharge data at both stations, 194 the ratios were linearly interpolated. Using the resulting gapless time-series of ratios as 195 estimators, the gap-filling of the discharge values was performed. In a similar way, the gaps 196 at WU17 were filled based on ratios to discharge values from WU14. 197

Figure 2 shows the gap-filled precipitation record as well as the gap-filled dischargehydrographs at WU14 and WU17 for the entire observation period from 2010 to 2024.









Figure 2: Precipitation and discharge time series at WU14 and WU17 (top and centre panels). Nitrate concentrations from weekly grab and event samples as well as the continuous nitrate concentrations measured with the TriOS sensor at WU14 (lower panel).

The average discharge during the 16-year record of WU14 was 9.61 l/sec (2.16 mm/day) and 1.47 l/sec (1.16) mm/day for the record of WU17 from August 2011 to the end of 2024 (Figure 2). The highest measured discharge was 50.57 and 28.95 mm/hour at WU14 and WU17 in the summer of 2021 (Figure 2), when a flood disaster occurred in western Germany (Saadi et al., 2023). Typically, the largest discharges are recorded in the winter half-year, while the summer periods are characterized by low discharge conditions, especially at WU17 (Figure 2).





213 *3.2 Physical parameters of water quality*

Several physical parameters of water quality (i.e. electrical conductivity, pH, redox 214 potential and temperature) are measured at a resolution of 10 minutes using a multi-probe 215 sensor (YSI 6820, YSI Inc.) at the stations WU14 and WU17. These measurements allow 216 the detection of short-term changes in the chemistry of stream water, e.g. after intense 217 218 precipitation events. A comparison of the statistical distributions of the physical parameter values (see Figures A1 and A2) illustrate that although the pH (median = 6.25 and 6.15 at 219 WU14 and WU17, respectively) and the redox potential (median = 259 mV and 287 mV at 220 221 WU14 and WU17, respectively) show a high degree of similarity, there is a significant difference in their electrical conductivities (median = $230 \ \mu\text{S/cm}$ and $60 \ \mu\text{S/cm}$ at WU14 222 and WU17, respectively). This difference can be explained by the heavy use of deicing salts 223 on the road that runs south through the Wüstebach catchment area (Płaczkowska et al., 224 2024). Furthermore, the mean stream water temperature at WU14 (9.4 °C) is slightly higher 225 than at WU17 (8.5 °C), reflecting the lower shading effect in the deforested Wüstebach 226 catchment (Figures A1 and A2). However, these differences should not limit the usability 227 of the reference stream for the investigation of the effects of deforestation on matter fluxes 228 in the Wüstebach catchment. 229

230

3.3 Solute concentrations

In order to determine the solute fluxes, both weekly sampling data of the flowing water (grab samples) and event-based samplings at stations WU14 and WU17 were employed. Event-driven sampling was carried out using autosamplers (AWS 2002, Ecotech, Germany) with an hourly sampling interval to detect rapid changes in water chemistry during high flow. Additionally, WU14 was equipped with an optical sensor (TriOS proPS, Rastede, Germany) to quantify alterations in nitrate concentrations at a 10-minute interval (Bogena et al., 2018).

For the sampling HDPE bottles were used, which were pre-rinsed with flowing water and
then completely filled to avoid air space. All samples were filtered in the laboratory (0.45
µm) and stored at 4°C prior to analysis. Major anions (Cl⁻, NO³⁻, SO₄⁻, NH₄⁺, PO₄³⁻) and
cations (Al³⁺, Fe_{tot}, Ca²⁺, Mg²⁺, Na⁺, K⁺) were measured using ion chromatography and
inductively coupled plasma–optical emission spectrometry (ICP-OES), respectively. The
DOC concentration was determined as non-purgeable organic carbon (Shimadzu TOC-





- VCPN). Given that the concentrations of NH4⁺ and PO4³⁻ were typically near or below the
 detection limit (0.06 and 0.08 mg/L), they were excluded from the data analysis.
 Figure 3 presents solute concentration boxplots of the most important macro- and
 microscopic discopled a burging and POC, determined from the most here also and
- 248 micronutrients, as well as dissolved aluminum and DOC, determined from the weekly and249 event-based stream water samples taken at WU14.
- 250



251



256

257 The solute concentrations in the Wüstebach stream were mostly low (median of all solutes = 3.52 mg/L), with Na⁺ showing the highest concentrations among the cations (median = 258 259 25.7 mg/L) and Cl⁻ as the dominant anion (median = 53.0 mg/L). The high NaCl content is 260 mainly caused by the large amount of road deicing salts entering the headwater of the Wüstebach catchment (Płaczkowska et al., 2024). In contrast, since the reference catchment 261 is not affected by these salt inputs (median $Na^+ = 2.90 \text{ mg/L}$, median $Cl^- = 4.94 \text{ mg/L}$, Figure 262 263 A3), the median of all solute concentrations in the reference stream was significantly lower 264 at 2.50 mg/L.

Using the nitrate and DOC concentration data, we selected the most appropriate method for calculating the hourly concentrations of all the other dissolved substances considered, which





267 are needed for the solute flux estimation (see section 3.4). In this way, the selected model should encompass both solutes that are more associated with subsurface transport processes 268 269 (i.e. nitrate) as well as overland flow transport processes (i.e. DOC). Mean nitrate 270 concentrations in both streams were generally low at 4.30 mg/L and 3.78 mg/L at WU14 and WU17, respectively. Overall, the measured stream nitrate concentrations during the 271 study period ranged from 0.64 to 14.65 mg/L at WU14 and 0.03 to 9.51 mg/L at WU17 (see 272 Figures 3 and A3) with strong seasonal variations (Figure 2). In addition, at WU14 the 273 nitrate concentration shows a long-term downward trend from 2016 onwards (Figure 2), 274 275 which is linked to the growth of new vegetation in the clear-cut area, which, compared to 276 the former spruce forest, takes up higher amounts of nitrogen from the soil (Płaczkowska et 277 al., 2022).

278 Regarding the export of dissolved organic matter, the behavior of both streams was very similar: DOC concentrations ranged at WU14 from 0.08 to 19.4 mg/L and at WU17 from 279 280 0.13 to 19.2 mg/L (see Figures 3 and A3). DOC concentrations were positively related to discharge only during the summer half year (Figure 4). This suggests that surface runoff 281 after high intensity precipitation events plays a crucial role in transporting DOC from soils 282 283 to surface waters (Strohmeier et al., 2013). The medium DOC concentration at WU14 (i.e. 284 4.82 mg/L) was about a quarter higher than at WU17 (i.e. 3.63 mg/L), which can be explained by the clear-cut measure that induced higher DOC concentrations over a few 285 years after the clear-cut (Robinson et al., 2022). 286

In contrast to DOC, nitrate concentrations show a strong positive trend with discharge during the winter months, while in the summer months this relationship was weaker, especially at WU14 (Figure 4). The non-linear correlations of DOC and nitrate concentrations with discharge present varying challenges for the solute concentration models, which is the reason for choosing these two solutes for the selection of the most appropriate method.







- Summer half-year - Winter half-year

294

Figure 4: Scatter diagrams of DOC and nitrate against the logarithmic discharge measured at stations WU14 and WU17 and differentiated according to summer and winter half-year. The LOESS (Locally Estimated Scatterplot Smoothing) trend lines, i.e. non-parametric curves that smooth data points by fitting local weighted regressions, indicate non-linearity, especially for DOC during the summer half-year.

300

301 *3.4 Solute flux calculation*

The measured weekly grab sample-based and event sampling-based stream solute 302 concentrations and associated discharge values from the discharge stations were employed 303 304 to generate models for the prediction of continuous hourly concentrations utilizing the R package LOADFLEX (Appling et al., 2015). Furthermore, LOADFLEX was employed to 305 construct composite models that rectify the inherent biases in the regression analysis 306 through the interpolation of prediction residuals and their subsequent incorporation into the 307 predicted values (Aulenbach and Hooper, 2006; Huntington et al., 1994). The optimal solute 308 309 concentration models were selected by minimizing the Akaike Information Criteria values through the utilization of the intrinsic functions provided by LOADFLEX. 310

Continuous solute concentration data are required to calculate hourly flux data. In order to find the most appropriate solute model for predicting continuous concentration observations, we compare the four basic types of models implemented in LOADFLEX (i.e. interpolation models, simple linear regression models, RLOADEST regression models and composite models). We use the nitrate concentrations measured at station WU14 as a reference example, as we also have high-resolution nitrate data from the TriOS probe that can be used for model validation.



The first and simplest model used in this study is the interpolation method, which has been used in many studies of solute and sediment fluxes (e.g. Buso et al. 2000, Vanni et al. 2001) when regression models are unsatisfactory (e.g. due to a weak relationship between the predictor and solute observations). Here, we employ the rectangular interpolation implemented in LOADFLEX, which is mathematically equivalent to a period-weighted averaging method (Likens, 2013).

The second solute concentration model used in this study is the linear regression method, which has long been an alternative to interpolation models for estimating substance flows, since it requires only limited time series of solute concentrations, as long as the predictor data, e.g. runoff, is available for the entire period of interest (Preston et al., 1989; Robertson and Roerish, 1999).

Next, we employed one of the more intricate RLOADEST regression models (Lorenz et al.,
2015; Runkel and De Cicco, 2017), specifically the seven-parameter model proposed by
Cohn et al. (1992), which is also incorporated in LOADFLEX:

332
$$ln(C) = \beta_0 + \beta_1 ln\left(\frac{Q}{\bar{Q}}\right) + \beta_2 \left[ln\left(\frac{Q}{\bar{Q}}\right)\right]^2 + \beta_3 \left(T - \tilde{T}\right) + \beta_4 \left(T - \tilde{T}\right) + \beta_5 sin(2\pi T) + \beta_6 cos(2\pi T) + \varepsilon$$
(2)

334 where C is the solute concentration, Q is the discharge, T is time in years, and \tilde{Q} and \tilde{T} are centering variables for which the average of the original variables is subtracted from each 335 individual value. The errors, denoted ε , are assumed to be independent and normally 336 distributed. The seven adjustment parameters, β_i , have to be fitted with the observation data. 337 The β_t parameters are used to scale the influences of O and T, as well as annual seasonality 338 on solute concentrations. Potential biases due to the logarithmic retransformation are 339 circumvented through an adjusted maximum likelihood estimation (AMLE) algorithm 340 (Appling et al., 2015). It should also be noted that the RLOADEST provides in total nine 341 342 different prediction models of which the seven-parameter model is the most complex (Kelly et al., 2018). We refer to Cohn et al. (1992) for more information on the model derivation 343 and application. 344

Finally, we employed the composite model suggested by Aulenbach and Hooper (2006) as a fourth method, which is mostly used in cases where regression predictions are available but exhibit medium-range biases (Appling et al., 2015). The composite method corrects systematic biases in solute concentrations predicted by a regression model by adjusting these predictions to observations with a linearly interpolated function of the model residuals

between the measurement time steps. Thus, by making use of all the available data, it should provide more accurate predictions at the temporal resolution of the predictors (Kelly et al., 2018). The composite model can use any regression model, which in this case is the RLOADEST regression model shown in Eq. (2). The residuals (e.g. absolute differences) are then calculated from the predictions and observations. The composite model then corrects the regression model predictions by adding the residuals to predictions or multiplying the predictions by the relative residuals.

357

363

373

358 *3.5 Statistical methods for model validation*

To validate the model results, three well-known statistical methods are used, namely, root mean square error (RMSE), Nash-Sutcliffe efficiency (NSE), and percent bias (PBIAS), which are briefly described here. RMSE is one of the most used error index statistics and is defined as follows:

$$RMSE = \sqrt{\frac{1}{n} \sum_{n=1}^{n} \left(C_i - \widehat{C}_i \right)^2}$$
(3)

in which \hat{C}_i is the vector of predicted solute concentrations, n is the number of data points and C_i is the vector of observed concentrations.

The Nash-Sutcliffe efficiency (NSE) determines the relative magnitude of the residual variance compared to the variance of observations and is computed as follows (Nash and Sutcliffe, 1970):

369 $NSE = 1 - \left[\frac{\sum_{n=1}^{n} (C_i - \hat{C}_i)^2}{\sum_{n=1}^{n} (C_i - \hat{C}_{mean})^2}\right]$ (4)

in which \hat{C}_{mean} is the arithmetic mean of predicted concentrations.

Finally, PBIAS quantifies the mean tendency of the simulated solute concentrations toexceed or fall below the observations and is calculated as follows (Gupta et al., 1999):

$$PBIAS = \left[\frac{\sum_{n=1}^{n} (C_i - \hat{C}_i)^2 \cdot 100}{\sum_{n=1}^{n} (C_i)}\right]$$
(5)

According to Moriasi et al. (2007), the output of a model for the concentration of dissolved
substances can be considered satisfactory if NSE exceeds 0.50 and PBIAS is within ±70 %.
We used an average quality measure, AQM, derived from NSE, RMSE and PBIAS to select
the most appropriate model for flow calculation. AQM is defined as:

$$AQM = \frac{NSE^* + RMSE^* + PBIAS^*}{3} \tag{6}$$

379	where NSE*, RMSE* and PBIAS* indicate the standardized metrics. Since NSE already lies
380	between 0 and 1, NSE corresponds to NSE* when negative values are set to 0. The
381	standardized RMSE [*] (i.e. RMSE mapped on a scale from 0 to 1) is calculated with:

$$RMSE^* = 1 - \frac{RMSE}{RMSE_{max}} \tag{7}$$

383 where RMSE_{max} is the worst RMSE value. Correspondingly, is PBIAS^{*} calculated using:

382

$$PBIAS^* = 1 - \frac{|PBIAS|}{PBIAS_{max}} \tag{8}$$

385 4 Results

386 *4.1 Comparison of different solute prediction models*

The panels in Figure 5 illustrate the nitrate prediction results at WU14 obtained using rectangular interpolation, linear regression, RLOADEST regression, and composite modelling for the whole observation period. The corresponding results at WU17 are presented in Figure A4.

391

Figure 5: Nitrate prediction models fitted to the measured values from weekly samples at WU14 using functions implemented in the LOADFLEX package. Subplots show observed concentrations (black dots) and predicted concentrations (orange lines) for 2010 to 2024 for rectangular interpolation (no3_interp), linear regression model (no3_lm), RLOADEST regression model (no3 reg2), and composite model (no3 comp).

In addition, Figure A5 shows both observed nitrate concentrations by the TriOS sensor andpredicted nitrate concentrations at WU14 for the years 2014 and 2020 to highlight the

399 performance differences between the various models. These two years were chosen because the data from the TriOS sensor is almost complete and because the nitrate concentrations 400 401 show significant differences. The discrepancies between the model performances are 402 particularly evident during the nitrate concentration peaks in July and December 2014, during which the linear and RLOADEST regression models severely underestimate nitrate 403 concentration, whereas the other methods perform well. The performance of the linear 404 regression model also deteriorates from 2020 onwards, leading to significantly excessive 405 406 nitrate concentrations, particularly in the winter months.

407 Since there is a general positive correlation between nitrate concentrations and runoff (Figure 4), the linear regression model always calculates increasing concentrations with 408 increasing runoff. However, there are obviously many inverted concentration peaks in 2020 409 (Figure A5), indicating dilution effects that were less pronounced in 2014 when nitrate 410 levels were generally higher due to clear-cutting. These observations are also reflected in 411 the efficiency statistics (Table 1), where in particular the linear regression model shows 412 poor quality indices for all variants. Overall, the rectangular interpolation and the composite 413 model showed the best performances. 414

415

Table 1: Summary statistics of the model validation using the hourly nitrate concentration
measured with the TriOS probe for the years 2014 and 2020 (green fields indicate the
highest quality, while red fields indicate the lowest quality).

Prediction model	Year	RMSE	NSE	PBIAS
Rectangular interpolation		1.18	0.33	9.85
Linear regression	2014	1.79	-0.54	28.79
RLOADEST (7 parameter)		1.77	-0.49	22.03
Composite model		1.11	0.41	10.32
Rectangular interpolation		0.40	0.74	3.34
Linear regression	2020	0.86	-0.22	-22.09
RLOADEST (7 parameter)		0.36	0.79	-2.64
Composite model		0.38	0.76	3.02

420	Table 2 summarizes the AQM values, which show that the composite model provides the
421	best model performance in predicting nitrate concentrations for the Wüstebach for the years
422	2014 and 2020.

423

Table 2: Average quality measures (AQM) for the four different solute prediction models(green fields indicate the highest quality, while red fields indicate the lowest quality).

Prediction model	AQM
Rectangular interpolation	0.62
Linear regression	0.13
RLOADEST (7 parameter)	0.46
Composite model	0.65

426

In a next step, the annual nitrate fluxes, which were determined on the basis of the 427 continuous nitrate concentrations from the various solution models, were compared 428 (Figure 6). As expected, the annual nitrate fluxes based on the linear model deviate 429 significantly from the other methods, which show a very similar trend. In addition, the 430 431 annual nitrate fluxes were also calculated on the basis of the weekly and event-based sampling dates (hereafter referred to only as "weekly sampling") and are also presented in 432 433 Figure 6 to see how the choice of the solute concentration model affects the calculation of the annual flux. In the dry year of 2018, the four models overestimate the flux derived from 434 weekly sampling by approx. 211 kg/year. Total nitrate fluxes for 2014 and 2020, based on 435 436 the continuous discharge and nitrate data from the TriOS sensor, were 2560.1 and 1143.3 kg/year, respectively. The linear model deviates the most from these values with 29.7 %, 437 whereas the rectangular interpolation and composite model deviate by only 2.94 % and 438 4.09 %, respectively. 439

440

Figure 6: Predictions of the annual nitrate fluxes at WU14 using different solute models,
weekly sampling data, as well as hourly nitrate concentration measurements with the TriOSsensor for 2014 and 2020.

444

The predicted DOC concentrations at WU14 and WU 17 are presented in Figure A6. To
assess the effects of the different concentration models on the calculation of the DOC flux,
annual DOC fluxes over the entire period were also calculated (Figure 7).

448

Figure 7: Predictions of the annual DOC fluxes at WU14 using different solute models aswell as the weekly sampling data.

451

452 Compared to the annual DOC flux calculated using the composite model, the linear
453 regression (42.1 %) and the RLOADEST method (48.5 %) show the greatest deviations,
454 while the 'weekly sampling' variant only deviates by 6 %. The smallest deviation in annual
455 DOC flux occurs in relation to rectangular interpolation methods is (-1.5%).

456 *4.2 Predicted solute concentrations*

- 457 Figures 8 and A7 show the predicted hourly concentrations of all considered solutes using
- the composite model at WU14 and WU17, respectively.

460 Figure 8: Predicted hourly concentrations of all considered solutes at WU14 using the461 composite model.

The solute concentrations at WU14 and WU17 exhibit sub-daily and seasonal variations of
varying intensity, which are obviously very well reproduced by the composite model.
However, it should be noted that the model produces a few concentration peaks for Al³⁺,
Ca²⁺, NO₃⁻ and DOC, which are not supported by the observations.

466 The annual average concentrations of the solutes considered show varying long-term trends (Figure A8a). For example, in the Wüstebach stream (WU14), the concentrations of Mg²⁺ 467 and NO₃⁻ show significant decreasing trends during the complete observation period, while 468 DOC and Fe^{2+} show significant increasing trends at a significance level of 0.05 (Table A3). 469 Interestingly, the reference stream (WU17) shows similar trends for most solutes, indicating 470 that the deforestation in the Wüstebach catchment is not the primary cause of these trends 471 (Figure A8b, Table A3). For example, the decreasing trend in nitrate concentration is 472 presumably related to the underplanting of beeches throughout the entire spruce stand, 473 which absorb additional nitrogen from the soil. The background to the underplanting is that 474 the entire spruce stand will be cleared by the mid-2030s, allowing a near-natural mixed 475 beech forest to develop. 476

An interesting anomaly is that in the Wüstebach the Fe^{2+} concentrations in 2024 were 477 significantly higher than in previous years (average Fe²⁺ concentration for 2010-2023 was 478 0.15 mg/L and for 2024 it was 0.57 mg/L, an increase of 283%). One hypothesis for the 479 high iron concentrations in the Wüstebach stream is that they are related to the anomalous 480 redox conditions in the soils of the riparian zone (Tittel et al., 2022). In support of this thesis 481 the alluvial soils in the riparian zone of the Wüstebach catchment (corresponds to the blue 482 colors in Figure 1) were saturated for longer periods than usual due to the wetter conditions 483 in the last two years (i.e. the mean annual precipitation for 2010-2022 was 1141 mm and 484 for 2023-2024 it was 1488 mm). Under waterlogged conditions, oxygen availability 485 decreases, creating an anaerobic environment that promotes microbial reduction of ferric 486 (Fe^{3+}) oxides to soluble ferrous (Fe^{2+}) iron (Ekström et al., 2016). If this occurs in 487 hydraulically connected soils, Fe^{2+} is transported into the stream. This hypothesis is also 488 489 supported by the exceptionally high Mn^{2+} concentrations in 2024 (average Mn^{2+} concentration for 2010-2023 was 0.09 mg/L and for 2024 it was 0.25 mg/L, an increase of 490 172%), since manganese (Mn⁴⁺) oxides are also reduced under waterlogged conditions 491 (Canfield et al., 1993). In addition, higher Fe²⁺ and Mn²⁺ concentrations were also found in 492 groundwater samples taken in the riparian zone in 2024 (data not shown), whereas Fe²⁺ and 493

- 494 Mn²⁺ concentrations in the reference stream (WU17) showed no increase in 2024 due to the lack of a larger riparian zone. 495 This example illustrates the advantage of long-term data series on the one hand, and on the 496 other hand, the existence of complementary data, such as the additional measurements in 497 the reference catchment, to reveal the complex relationships in solute transport. 498 499 4.2 Predicted solute fluxes 500 501 Figure 9 shows the predicted annual fluxes per unit area (km²) of all considered solutes using the composite model for the Wüstebach and reference streams, respectively. The 502 conversion to a unit area facilitates the comparison of fluxes for different streams. In the 503 reference stream, the annual solute fluxes of the ions considered ranged from less than 39.1 504 kg km⁻² a^{-1} (Fe²⁺) to 4929.5 kg km⁻² a^{-1} (SO⁴⁻). The fluxes of NO₃⁻ and CL⁻ are in the middle 505 range at 1958.2 and 2033.7 kg km⁻² a⁻¹, respectively. The cations with the highest mean 506 annual fluxes were Ca and Na with 1128.6 and 1127.1 kg km⁻² a⁻¹, respectively. The lowest 507 annual solute fluxes occurred in the drought year 2018 (47% of the average total annual 508 solute fluxes), mainly due to low runoff in that year (67% of the long-time average observed 509 510 runoff at WU17).
- In the Wüstebach stream, the average annual total solute flux (i.e. 77.9 t km⁻² a⁻¹) was six times higher than in the reference stream (i.e. 13.0 t km⁻² a⁻¹). One reason for this was that the total annual flow of dissolved substances of Na⁺ and Cl⁻ was sixteen and eighteen times higher in the Wüstebach stream, i.e. 18.6 and 37.4 t km⁻² a⁻¹, respectively, due to the input of de-icing salt on the road that crosses the catchment (Płaczkowska et al., 2022, 2024).
 In the Wüstebach stream, five solutes show significant long-term trends in the annual fluxes:
- Fe²⁺, K⁺ and DOC show significant increasing trends while those of NO₃⁻ and SO₄²⁻ show
 significant decreasing trends (Figure A9a, Table A3). In contrast, the reference stream
 shows only significant decreasing trends of NO₃⁻ (Figure A9b, Table A3).
- 520

522

Figure 9: Annual fluxes of all considered solutes using the composite model for theWüstebach (WU14) and reference stream (WU17), respectively.

526 5 Conclusions and outlook

527 In this data paper, we present 16 years of hourly stream-water flux data of nine continuously 528 monitored macro- and micronutrients, as well as dissolved ionic aluminum and DOC, in the 529 Wüstebach catchment. The basis for the solute fluxes is continuous discharge and water 530 quality data collected as part of a deforestation experiment. To enable analysis of the effects of the deforestation measure on solute transport, these data were also collected in a 531 532 neighboring catchment (i.e. reference stream). To ascertain the most appropriate method for predicting continuous solute concentrations, four well-known methods were applied and 533 534 compared with high-resolution nitrate measurements taken in the Wüstebach stream over a 535 period of two years. Preliminary data analysis of the comprehensive data set on concentrations and stream-water 536

fluxes of dissolved substances shows great potential for analyzing the influence of deforestation and afforestation on the fluxes of dissolved substances at the catchment scale. In addition, the effects of extreme dry and wet years on the concentrations and fluxes of substances can be investigated in more detail using the environmental monitoring data available at the Wüstebach site.

The data processing employed in this study is currently state of the art. However, this may be subject to alteration considering future research. Consequently, we offer raw data and undertake to revise the published dataset with an incremental version, should novel data processing methodologies be endorsed in the future. Since the measurements in Wüstebach are still ongoing, we also plan to update the stream-water flux time series regularly, e.g. every three years.

548

550 Data access and availability

- The described long-term data set, the R scripts to generate the plots from the data and the 551 geospatial data presented in Figure 1 are directly accessible via a digital object identifier 552 (DOI) and is freely available at https://doi.org/10.26165/JUELICH-DATA/AKAMNQ 553 554 (Bogena and Herrmann, 2025). All timestamps associated with the data records are recorded 555 in Coordinated Universal Time (UTC). Table A1 in the Appendix provides an overview of 556 the published data files. Further data from instrumentation in the Wüstebach catchment are freely available via the TERENO data portal TEODOOR (http://teodoor.icg.kfa-557 558 juelich.de/). In addition, inquiries about data and research collaborations can be directed to the lead author (h.bogena@fz-juelich.de). 559
- 560

561 Competing interests

562 The contact author has declared that neither they nor their co-authors have any competing 563 interests.

564

565 Author contributions

HRB conceptualized the study, led the data compilation and editing of the manuscript. FH
and HRB developed the R scripts for data pre-processing, solute estimation post-processing
and performed the formal analysis. HRB wrote the first manuscript draft, and all authors
contributed to editing the final version of the manuscript.

570

571 Acknowledgements

572 The Agrosphere Institute of the Forschungszentrum Jülich GmbH funds the operation of the Wüstebach catchment, including instrumentation maintenance and salaries of all field and 573 laboratory technicians. These data were collected using support by TERENO (Terrestrial 574 575 Environmental Observatories) funded by the Helmholtz-Gemeinschaft and can be used without any further approval or charges. For the use of the data in a publication, presentation 576 or other research material, please quote the data according to the TERENO disclaimer. We 577 578 thank Antonio Voss, Phillip Meulendick, Ferdinand Engels, Rainer Harms, Werner Küpper, 579 Andrea Ecker, Stephanie Stork, Sirgit Kummer, Herbert Rützel, Martina Krause, and Holger Wissel for supporting the sampling, chemical analysis and regular maintenance of 580 581 the experimental set-up. Finally, we would like to thank the Eifel National Park for its 582 constant support of our research.

583 References

584Appling, A. P., Leon, M. C., and McDowell, W. H.: Reducing bias and quantifying uncertainty in585watershed flux estimates: the R package loadflex, Ecosphere, 6, 1–25,586https://doi.org/10.1890/ES14-00517.1, 2015.

587 Aulenbach, B. T. and Hooper, R. P.: The composite method: an improved method for stream-water 588 solute load estimation, Hydrol. Process., 20, 3029–3047, https://doi.org/10.1002/hyp.6147, 2006.

589 Birgand, F., Faucheux, C., Gruau, G., Augeard, B., Moatar, F., and Bordenave, P.: Uncertainties in 590 Assessing Annual Nitrate Loads and Concentration Indicators: Part 1. Impact of Sampling 591 Algorithms, Frequency and Load Estimation Trans. ASABE, 53, 437-446, 592 https://doi.org/10.13031/2013.29584, 2010.

Bogena, H. and Herrmann, F.: Wüstebach data paper: Long-term hourly solute flux data 2010-2024,
 https://doi.org/10.26165/JUELICH-DATA/AKAMNQ, 2025.

Bogena, H. R., Herbst, M., Huisman, J. A., Rosenbaum, U., Weuthen, A., and Vereecken, H.:
Potential of Wireless Sensor Networks for Measuring Soil Water Content Variability, Vadose Zone
J., 9, 1002–1013, https://doi.org/10.2136/vzj2009.0173, 2010.

Bogena, H. R., Bol, R., Borchard, N., Brüggemann, N., Diekkrüger, B., Drüe, C., Groh, J., Gottselig,
N., Huisman, J. A., Lücke, A., Missong, A., Neuwirth, B., Pütz, T., Schmidt, M., Stockinger, M., Tappe,
W., Weihermüller, L., Wiekenkamp, I., and Vereecken, H.: A terrestrial observatory approach to
the integrated investigation of the effects of deforestation on water, energy, and matter fluxes,
Sci. China Earth Sci., 58, 61–75, https://doi.org/10.1007/s11430-014-4911-7, 2015.

Bogena, H. R., Montzka, C., Huisman, J. A., Graf, A., Schmidt, M., Stockinger, M., von Hebel, C.,
Hendricks-Franssen, H. J., van der Kruk, J., Tappe, W., Lücke, A., Baatz, R., Bol, R., Groh, J., Pütz, T.,
Jakobi, J., Kunkel, R., Sorg, J., and Vereecken, H.: The TERENO-Rur Hydrological Observatory: A
Multiscale Multi-Compartment Research Platform for the Advancement of Hydrological Science,
Vadose Zone J., 17, 180055, https://doi.org/10.2136/vzj2018.03.0055, 2018.

Bogena, H. R., Stockinger, M. P., and Lücke, A.: Long-term stable water isotope and runoff data for
 the investigation of deforestation effects on the hydrological system of the Wüstebach catchment,
 Germany, Hydrol. Process., 35, e14006, https://doi.org/10.1002/hyp.14006, 2021.

Bol, R., Lücke, A., Tappe, W., Kummer, S., Krause, M., Weigand, S., Pütz, T., and Vereecken, H.:
Spatio-temporal Variations of Dissolved Organic Matter in a German Forested Mountainous
Headwater Catchment, Vadose Zone J., 14, vzj2015.01.0005,
https://doi.org/10.2136/vzj2015.01.0005, 2015.

Brunet, C. E., Gemrich, E. R. C., Biedermann, S., Jacobson, P. J., Schilling, K. E., Jones, C. S., and
Graham, A. M.: Nutrient capture in an Iowa farm pond: Insights from high-frequency observations,
J. Environ. Manage., 299, 113647, https://doi.org/10.1016/j.jenvman.2021.113647, 2021.

618 Canfield, D. E., Thamdrup, B., and Hansen, J. W.: The anaerobic degradation of organic matter in
619 Danish coastal sediments: Iron reduction, manganese reduction, and sulfate reduction, Geochim.
620 Cosmochim. Acta, 57, 3867–3883, https://doi.org/10.1016/0016-7037(93)90340-3, 1993.

621 Chen, M., Jomaa, S., Lausch, A., Beudert, B., Ghaffar, S., Jia, W., and Rode, M.: Impact of Forest
622 Dieback on Hydrology and Nitrogen Export Using a New Dynamic Water Quality Model, Water

623 Resour. Res., 60, e2024WR037341, https://doi.org/10.1029/2024WR037341, 2024.

Coble, A. A., Wymore, A. S., Shattuck, M. D., Potter, J. D., and McDowell, W. H.: Multiyear Trends
in Solute Concentrations and Fluxes From a Suburban Watershed: Evaluating Effects of 100-Year
Flood Events, J. Geophys. Res. Biogeosciences, 123, 3072–3087,
https://doi.org/10.1029/2018JG004657, 2018.

628 Cohn, T. A., Caulder, D. L., Gilroy, E. J., Zynjuk, L. D., and Summers, R. M.: The validity of a simple 629 statistical model for estimating fluvial constituent loads: An Empirical study involving nutrient 630 loads entering Chesapeake Bay, Water Resour. Res., 28. 2353-2363, 631 https://doi.org/10.1029/92WR01008, 1992.

Cox, N. J., Warburton, J., Armstrong, A., and Holliday, V. J.: Fitting concentration and load rating
curves with generalized linear models, Earth Surf. Process. Landf., 33, 25–39,
https://doi.org/10.1002/esp.1523, 2008.

Deuffic, P., Garms, M., He, J., Brahic, E., Yang, H., and Mayer, M.: Forest Dieback, a Tangible Proof
of Climate Change? A Cross-Comparison of Forest Stakeholders' Perceptions and Strategies in the
Mountain Forests of Europe and China, Environ. Manage., 66, 858–872,
https://doi.org/10.1007/s00267-020-01363-9, 2020.

Ekström, S. M., Regnell, O., Reader, H. E., Nilsson, P. A., Löfgren, S., and Kritzberg, E. S.: Increasing
concentrations of iron in surface waters as a consequence of reducing conditions in the catchment
area, J. Geophys. Res. Biogeosciences, 121, 479–493, https://doi.org/10.1002/2015JG003141,
2016.

Feller, M. C.: Forest harvesting and streamwater inorganic chemistry in western north America: a
review, J. Am. Water Resour. Assoc., 41, 785–811, https://doi.org/10.1111/j.17521688.2005.tb04464.x, 2005.

Georgiev, K. B., Beudert, B., Bässler, C., Feldhaar, H., Heibl, C., Karasch, P., Müller, J., Perlík, M.,
Weiss, I., and Thorn, S.: Forest disturbance and salvage logging have neutral long-term effects on
drinking water quality but alter biodiversity, For. Ecol. Manag., 495, 119354,
https://doi.org/10.1016/j.foreco.2021.119354, 2021.

Gottselig, N., Wiekenkamp, I., Weihermüller, L., Brüggemann, N., Berns, A. E., Bogena, H. R.,
Borchard, N., Klumpp, E., Lücke, A., Missong, A., Pütz, T., Vereecken, H., Huisman, J. A., and Bol, R.:
A Three-Dimensional View on Soil Biogeochemistry: A Dataset for a Forested Headwater
Catchment, J. Environ. Qual., 46, 210–218, https://doi.org/10.2134/jeq2016.07.0276, 2017.

Graf, A., Bogena, H. R., Drüe, C., Hardelauf, H., Pütz, T., Heinemann, G., and Vereecken, H.:
Spatiotemporal relations between water budget components and soil water content in a forested
tributary catchment, Water Resour. Res., 50, 4837–4857,
https://doi.org/10.1002/2013WR014516, 2014a.

Graf, A., Bogena, H. R., Drüe, C., Hardelauf, H., Pütz, T., Heinemann, G., and Vereecken, H.:
Spatiotemporal relations between water budget components and soil water content in a forested
tributary catchment, Water Resour. Res., 50, 4837–4857,
https://doi.org/10.1002/2013WR014516, 2014b.

Gupta, H. V., Sorooshian, S., and Yapo, P. O.: Status of Automatic Calibration for Hydrologic Models:
Comparison with Multilevel Expert Calibration, J. Hydrol. Eng., 4, 135–143, https://doi.org/10.1061/(ASCE)1084-0699(1999)4:2(135), 1999.

Harley, J. R., Biles, F. E., Brooks, M. K., Fellman, J., Hood, E., and D'Amore, D. V.: Riverine Dissolved
Inorganic Carbon Export From the Southeast Alaskan Drainage Basin With Implications for Coastal
Ocean Processes, J. Geophys. Res. Biogeosciences, 128, e2023JG007609,
https://doi.org/10.1029/2023JG007609, 2023.

Heistermann, M., Bogena, H., Francke, T., Güntner, A., Jakobi, J., Rasche, D., Schrön, M., Döpper,
V., Fersch, B., Groh, J., Patil, A., Pütz, T., Reich, M., Zacharias, S., Zengerle, C., and Oswald, S.: Soil
moisture observation in a forested headwater catchment: combining a dense cosmic-ray neutron
sensor network with roving and hydrogravimetry at the TERENO site Wüstebach, Earth Syst. Sci.
Data, 14, 2501–2519, https://doi.org/10.5194/essd-14-2501-2022, 2022.

- Hirsch, R. M. and De Cicco, L. A.: User guide to exploration and graphics for RivEr Trends (EGRET)
 and data Retrieval: R packages for hydrologic data. Version 2.0., U.S. Geological Survey, Reston,
 Virginia, USA, 2015.
- Huntington, T. G., Hooper, R. P., and Aulenbach, B. T.: Hydrologic processes controlling sulfate
 mobility in a small forested watershed, Water Resour. Res., 30, 283–295,
 https://doi.org/10.1029/93WR02950, 1994.
- Keller, H. M.: Der Chemismus kleiner Bache in teilweise bewaldeten Einzugsgebieten in der
 Flyschzone eines Voralpentales, Mitteilgn Schweiz. Anst. F Forstl. Vers., 46, 113–155, 1970.
- Kelly, P. T., Vanni, M. J., and Renwick, W. H.: Assessing uncertainty in annual nitrogen, phosphorus,
 and suspended sediment load estimates in three agricultural streams using a 21-year dataset,
 Environ. Monit. Assess., 190, 91, https://doi.org/10.1007/s10661-018-6470-4, 2018.
- Kong, X., Ghaffar, S., Determann, M., Friese, K., Jomaa, S., Mi, C., Shatwell, T., Rinke, K., and Rode,
 M.: Reservoir water quality deterioration due to deforestation emphasizes the indirect effects of
 global change, Water Res., 221, 118721, https://doi.org/10.1016/j.watres.2022.118721, 2022.
- Likens, G. E.: Biogeochemistry of a Forested Ecosystem, Springer, New York, NY,
 https://doi.org/10.1007/978-1-4614-7810-2, 2013.
- Likens, G. E., Bormann, F. H., Johnson, N. M., Fisher, D. W., and Pierce, R. S.: Effects of Forest
 Cutting and Herbicide Treatment on Nutrient Budgets in the Hubbard Brook WatershedEcosystem, Ecol. Monogr., 40, 23–47, https://doi.org/10.2307/1942440, 1970.
- Lorenz, D. R., Runkel, R., and De Cicco, L.: Rloadest: river load estimation, U.S. Geol. Surv. Software,
 Mounds View, Minnesota, USA, 2015.
- McDowell, W. H., McDowell, W. G., Potter, J. D., and Ramírez, A.: Nutrient export and elemental
 stoichiometry in an urban tropical river, Ecol. Appl., 29, e01839,
 https://doi.org/10.1002/eap.1839, 2019.
- Mikkelson, K. M., Dickenson, E. R. V., Maxwell, R. M., McCray, J. E., and Sharp, J. O.: Water-quality
 impacts from climate-induced forest die-off, Nat. Clim. Change, 3, 218–222,
 https://doi.org/10.1038/nclimate1724, 2013.
- Mollenhauer, H., Kasner, M., Haase, P., Peterseil, J., Wohner, C., Frenzel, M., Mirtl, M., Schima, R.,
 Bumberger, J., and Zacharias, S.: Long-term environmental monitoring infrastructures in Europe:
 observations, measurements, scales, and socio-ecological representativeness, Sci. Total Environ.,
 624, 968–978, https://doi.org/10.1016/j.scitotenv.2017.12.095, 2018.

Moriasi, D. N., Arnold, J. G., Van Liew, M. W., Bingner, R. L., Harmel, R. D., and Veith, T. L.: Model
 Evaluation Guidelines for Systematic Quantification of Accuracy in Watershed Simulations, Trans.
 ASABE, 50, 885–900, https://doi.org/10.13031/2013.23153, 2007.

Musolff, A., Tarasova, L., Rinke, K., and Ledesma, J. L. J.: Forest Dieback Alters Nutrient Pathways
in a Temperate Headwater Catchment, Hydrol. Process., 38, e15308,
https://doi.org/10.1002/hyp.15308, 2024.

Nash, J. E. and Sutcliffe, J. V.: River flow forecasting through conceptual models part I — A
discussion of principles, J. Hydrol., 10, 282–290, https://doi.org/10.1016/0022-1694(70)90255-6,
1970.

Ney, P., Graf, A., Bogena, H., Diekkrüger, B., Drüe, C., Esser, O., Heinemann, G., Klosterhalfen, A.,
Pick, K., Pütz, T., Schmidt, M., Valler, V., and Vereecken, H.: CO2 fluxes before and after partial
deforestation of a Central European spruce forest, Agric. For. Meteorol., 274, 61–74,
https://doi.org/10.1016/j.agrformet.2019.04.009, 2019.

Obladen, N., Dechering, P., Skiadaresis, G., Tegel, W., Keßler, J., Höllerl, S., Kaps, S., Hertel, M.,
Dulamsuren, C., Seifert, T., Hirsch, M., and Seim, A.: Tree mortality of European beech and Norway
spruce induced by 2018-2019 hot droughts in central Germany, Agric. For. Meteorol., 307, 108482,
https://doi.org/10.1016/j.agrformet.2021.108482, 2021.

Płaczkowska, E., Mostowik, K., Bogena, H. R., and Leuchner, M.: The Impact of Partial Deforestation
on Solute Fluxes and Stream Water Ionic Composition in a Headwater Catchment, Water, 15, 107,
https://doi.org/10.3390/w15010107, 2022.

Płaczkowska, E., Kijowska-Strugała, M., Ketzler, G., Bogena, H. R., and Leuchner, M.: Solute fluxes
in headwater catchments with contrasting anthropogenic impact, Geomorphology, 454, 109166,
https://doi.org/10.1016/j.geomorph.2024.109166, 2024a.

Płaczkowska, E., Kijowska-Strugała, M., Ketzler, G., Bogena, H. R., and Leuchner, M.: Solute fluxes
in headwater catchments with contrasting anthropogenic impact, Geomorphology, 454, 109166,
https://doi.org/10.1016/j.geomorph.2024.109166, 2024b.

Preston, S. D., Bierman, V. J., and Silliman, S. E.: An evaluation of methods for the estimation of
tributary mass loads, Water Resour. Res., 25, 1379–1389,
https://doi.org/10.1029/WR025i006p01379, 1989.

Rajwa-Kuligiewicz, A. and Bojarczuk, A.: Streamflow response to catastrophic windthrow and
forest recovery in subalpine spruce forest, J. Hydrol., 634, 131078,
https://doi.org/10.1016/j.jhydrol.2024.131078, 2024.

Robertson, D. M. and Roerish, E. D.: Influence of various water quality sampling strategies on load
estimates for small streams, Water Resour. Res., 35, 3747–3759,
https://doi.org/10.1029/1999WR900277, 1999.

Robinson, K.-L., Bogena, H. R., Wang, Q., Cammeraat, E., and Bol, R.: Effects of deforestation on
dissolved organic carbon and nitrate in catchment stream water revealed by wavelet analysis,
Front. Water, 4, 1003693, https://doi.org/10.3389/frwa.2022.1003693, 2022a.

Robinson, K.-L., Bogena, H. R., Wang, Q., Cammeraat, E., and Bol, R.: Effects of deforestation on
 dissolved organic carbon and nitrate in catchment stream water revealed by wavelet analysis,
 Front. Water, 4, 1003693, https://doi.org/10.3389/frwa.2022.1003693, 2022b.

- Rosenbaum, U., Bogena, H. R., Herbst, M., Huisman, J. A., Peterson, T. J., Weuthen, A., Western, A.
 W., and Vereecken, H.: Seasonal and event dynamics of spatial soil moisture patterns at the small
 catchment scale, Water Resour. Res., 48, 2011WR011518,
 https://doi.org/10.1029/2011WR011518, 2012.
- 750 Runkel, R. and De Cicco, L.: USGS-R/rloadest: River Load Estimation, 2017.
- Runkel, R. L., Crawford, C. G., and Timothy, A. C.: Load Estimator (LOADEST): A FORTRAN program
 for estimating constituent loads in streams and rivers, 2004.
- Saadi, M., Furusho-Percot, C., Belleflamme, A., Trömel, S., Kollet, S., and Reinoso-Rondinel, R.:
 Comparison of Three Radar-Based Precipitation Nowcasts for the Extreme July 2021 Flooding
 Event in Germany, J. Hydrometeorol., 24, 1241–1261, https://doi.org/10.1175/JHM-D-22-0121.1,
 2023.
- Stockinger, M. P., Bogena, H. R., Lücke, A., Diekkrüger, B., Weiler, M., and Vereecken, H.: Seasonal
 soil moisture patterns: Controlling transit time distributions in a forested headwater catchment,
 Water Resour. Res., 50, 5270–5289, https://doi.org/10.1002/2013WR014815, 2014.
- Strohmeier, S., Knorr, K.-H., Reichert, M., Frei, S., Fleckenstein, J. H., Peiffer, S., and Matzner, E.:
 Concentrations and fluxes of dissolved organic carbon in runoff from a forested catchment:
 insights from high frequency measurements, Biogeosciences, 10, 905–916,
 https://doi.org/10.5194/bg-10-905-2013, 2013.
- Tittel, J., Büttner, O., Friese, K., Lechtenfeld, O. J., Schuth, S., Von Tümpling, W., and Musolff, A.:
 Iron Exports From Catchments Are Constrained by Redox Status and Topography, Glob.
 Biogeochem. Cycles, 36, e2021GB007056, https://doi.org/10.1029/2021GB007056, 2022.
- Vilhar, U., Kermavnar, J., Kozamernik, E., Petrič, M., and Ravbar, N.: The effects of large-scale forest
 disturbances on hydrology An overview with special emphasis on karst aquifer systems, Earth Sci. Rev., 235, 104243, https://doi.org/10.1016/j.earscirev.2022.104243, 2022.
- Walker, W. W.: Simplified procedures for eutrophication assessment and prediction: user manual,
 U.S. Army Engineer Waterways Experiment Station, Vicksburg, Mississippi, USA, 1996.
- Wang, J., Bogena, H., Süß, T., Graf, A., Weuthen, A., and Brüggemann, N.: Investigating the controls
 on greenhouse gas emission in the riparian zone of a small headwater catchment using an
 automated monitoring system, Vadose Zone J., 20, https://doi.org/10.1002/vzj2.20149, 2021.
- Wang, J., Bouchez, J., Dolant, A., Floury, P., Stumpf, A. J., Bauer, E., Keefer, L., Gaillardet, J., Kumar,
 P., and Druhan, J. L.: Sampling frequency, load estimation and the disproportionate effect of
 storms on solute mass flux in rivers, Sci. Total Environ., 906, 167379,
 https://doi.org/10.1016/j.scitotenv.2023.167379, 2024.
- 779 Wang, Q., Qu, Y., Robinson, K.-L., Bogena, H., Graf, A., Vereecken, H., Tietema, A., and Bol, R.: 780 Deforestation alters dissolved organic carbon and sulfate dynamics in a mountainous headwater 781 For. Glob. 1044447, catchment—A wavelet analysis, Front. Change, 5, 782 https://doi.org/10.3389/ffgc.2022.1044447, 2022.
- Weigand, S., Bol, R., Reichert, B., Graf, A., Wiekenkamp, I., Stockinger, M., Luecke, A., Tappe, W.,
 Bogena, H., Puetz, T., Amelung, W., and Vereecken, H.: Spatiotemporal Analysis of Dissolved
 Organic Carbon and Nitrate in Waters of a Forested Catchment Using Wavelet Analysis, Vadose
 Zone J., 16, vzj2016.09.0077, https://doi.org/10.2136/vzj2016.09.0077, 2017.

- Wiekenkamp, I., Huisman, J. A., Bogena, H. R., Graf, A., Lin, H. S., Drüe, C., and Vereecken, H.:
 Changes in measured spatiotemporal patterns of hydrological response after partial deforestation
 in a headwater catchment, J. Hydrol., 542, 648–661,
 https://doi.org/10.1016/j.jhydrol.2016.09.037, 2016a.
- Wiekenkamp, I., Huisman, J. A., Bogena, H. R., Lin, H. S., and Vereecken, H.: Spatial and temporal
 occurrence of preferential flow in a forested headwater catchment, J. Hydrol., 534, 139–149,
 https://doi.org/10.1016/j.jhydrol.2015.12.050, 2016b.
- Wiekenkamp, I., Huisman, J. A., Bogena, H. R., and Vereecken, H.: Effects of Deforestation on Water
 Flow in the Vadose Zone, Water, 12, 35, https://doi.org/10.3390/w12010035, 2019.
- Worrall, F., Howden, N. J. K., and Burt, T. P.: Assessment of sample frequency bias and precision in
 fluvial flux calculations An improved low bias estimation method, J. Hydrol., 503, 101–110,
 https://doi.org/10.1016/j.jhydrol.2013.08.048, 2013.
- Xu, C., Förster, M., Beckschäfer, P., Talkner, U., Klinck, C., and Kleinschmit, B.: Modeling European
 beech defoliation at a regional scale gradient in Germany from northern lowlands to central
 uplands using geo-ecological parameters, Sentinel-2 and National Forest Condition Survey data,
 For. Ecol. Manag., 576, 122383, https://doi.org/10.1016/j.foreco.2024.122383, 2025.
- Zacharias, S., Bogena, H., Samaniego, L., Mauder, M., Fuß, R., Pütz, T., Frenzel, M., Schwank, M.,
 Baessler, C., Butterbach-Bahl, K., Bens, O., Borg, E., Brauer, A., Dietrich, P., Hajnsek, I., Helle, G.,
 Kiese, R., Kunstmann, H., Klotz, S., Munch, J. C., Papen, H., Priesack, E., Schmid, H. P., Steinbrecher,
 R., Rosenbaum, U., Teutsch, G., and Vereecken, H.: A Network of Terrestrial Environmental
 Observatories in Germany, Vadose Zone J., 10, 955–973, https://doi.org/10.2136/vzj2010.0139,
 2011.
- Zacharias, S., Loescher, H. W., Bogena, H., Kiese, R., Schrön, M., Attinger, S., Blume, T., Borchardt,
 D., Borg, E., Bumberger, J., Chwala, C., Dietrich, P., Fersch, B., Frenzel, M., Gaillardet, J., Groh, J.,
 Hajnsek, I., Itzerott, S., Kunkel, R., Kunstmann, H., Kunz, M., Liebner, S., Mirtl, M., Montzka, C.,
 Musolff, A., Pütz, T., Rebmann, C., Rinke, K., Rode, M., Sachs, T., Samaniego, L., Schmid, H. P., Vogel,
 H., Weber, U., Wollschläger, U., and Vereecken, H.: Fifteen Years of Integrated Terrestrial
 Environmental Observatories (TERENO) in Germany: Functions, Services, and Lessons Learned,
 Earths Future, 12, e2024EF004510, https://doi.org/10.1029/2024EF004510, 2024.
- 816

817 Appendix

818 Table A1: List and description of published data files.

File name	Short description
precipitation_combined.txt	Daily precipitation gap-filled and combined from two stations
readme.txt and readme.xlsx	Documentation of data tables and files, meaning of the headers, units, etc.
WU14_discharge.txt	Measured and gap-filled discharge record at station WU14
WU14_DOC_annual_flux.txt	Annual flux of DOC based on discontinuous sampling and the four solute prediction models at station WU14
WU14_DOC_concentration_modelling.txt	Hourly discharge and DOC concentrations modelled using the four solute prediction models at station WU14
WU14_elements_compositeModel_annual_flux.txt	Annual flux of elements at station WU14, modelled using the composite model
WU14_elements_compositeModel_annual_flux_per_sq uare_kilometer.txt	Annual flux of elements at station WU14, modelled using the composite model and converted to the catchment area
WU14_elements_compositeModel_concentrations_flux .txt	Hourly discharge, concentration and flux of elements at station WU14, modelled using the composite model
WU14_elements_ds_concentrations.txt	Discharge and concentration of elements from discontinuous sampling at station WU14
WU14_nitrate_measurements1.txt	Daily discharge and nitrate concentration from discontinuous sampling at station WU14
WU14_nitrate_measurements2_trios.txt	Hourly nitrate concentration measured with the TriOS sensor at station WU14
WU14_nitrate_annual_flux.txt	Annual flux of nitrate based on discontinuous sampling and the four solute prediction models at station WU14
WU14_nitrate_concentration_modelling.txt	Hourly discharge and nitrate concentration modelled using the four solute prediction models at station WU14
WU14_physicalParameters_ds.txt	Physical parameters from discontinuous sampling at station WU14
WU17_discharge.txt	Measured and gap-filled discharge record at station WU17
WU17_DOC_annual_flux.txt	Annual flux of DOC based on discontinuous sampling and the four solute prediction models at station WU17
WU17_DOC_concentration_modelling.txt	Hourly discharge and DOC concentrations modelled using the four solute prediction models at station WU17
WU17_elements_compositeModel_annual_flux.txt	Annual flux of elements at station WU17, modelled using the composite model
WU17_elements_compositeModel_annual_flux_per_sq uare_kilometer.txt	Annual flux of elements at station WU17, modelled using the composite model and converted to the catchment area
WU17_elements_compositeModel_concentrations_flux .txt	Hourly discharge, concentration and flux of elements at station WU17, modelled using the composite model
WU17_elements_ds_concentrations.txt	Discharge and concentration of elements from discontinuous sampling at station WU17
WU17_physicalParameters_ds.txt	Physical parameters from discontinuous sampling at station WU17
WU17_nitrate_concentration_modelling.txt	Hourly discharge and nitrate concentration modelled using the four solute prediction models at station WU17

820

- Table A2: Type of measurement and geographical coordinates of the TERENO stations
- 822 used in this study.

823

Station	Type of measurement	Latitude	Longitude
WU10	Discharge station	50.5038	6.3335
WU14	Discharge station	50.505	6.3334
WU17	Discharge station	50.5052	6.3341
WU_K_2	Meteorological station	50.5031	6.3360

- 825 Table A3: Results of the annual trend analysis. Significant trends (i.e. p-value < 0.05) are
 - Data Solute Intercept \mathbb{R}^2 Slope p value 0.000001 NO3 604.50 -0.297 0.851 DOC -324.19 0.163 0.398 0.0117 119.17 -0.057 0.325 0.0265 Mg Fe -31.40 0.016 0.306 0.0326 WU14 annual SO4 -114.87 0.061 0.226 0.0735 175.40 0.217 average Ca -0.084 0.0804concentration Κ -10.99 0.006 0.193 0.1016 Cl 1157.28 -0.547 0.118 0.2104 Mn -9.09 0.005 0.093 0.2704 Na 334.14 -0.153 0.065 0.3594 -0.58 0.000 0.004 Al 0.8188 NO3 507119.10 -249.129 0.649 0.0003 21.914 0.471 Κ -43559.88 0.0047DOC -325597.61 0.391 162.856 0.0127 SO4 -364719.16 184.203 0.348 0.0206 -20327.10 0.315 Fe 10.136 0.0296 WU14 annual flux 0.063 -420968.18 218.398 0.3651 Na Mn -4334.12 2.187 0.042 0.4655 Cl -549782.46 292.029 0.040 0.4756 Ca -9571.81 7.062 0.002 0.8744 4695.57 -0.961 0.0000.9724 Mg 0.000 Al 226.02 -0.062 0.9730 NO3 588.80 -0.290 0.853 0.000007 DOC -6.01 0.003 0.637 0.001169.69 -0.033 0.601 0.0019 Mg 43.41 -0.020 0.461 0.0107 Fe WU17 annual SO4 -20.87 0.011 0.446 0.0126 average Ca -49.82 0.026 0.376 0.0259 concentration -79.78 0.327 0.0412 Κ 0.041 Cl 0.001 0.200 0.1258 -1.01 0.001 0.195 Mn -1.95 0.1312 Na -57.18 0.031 0.120 0.2472 Al -11.25 0.011 0.035 0.5391 NO3 326338.35 -160.662 0.599 0.0019 -11016.39 0.1674 Κ 5.574 0.166 DOC 40419.10 -19.428 0.094 0.3071 SO4 28943.06 -13.813 0.060 0.4203 -583.46 0.299 0.045 0.4845 Fe AU17 annual flux Na -280.260.147 0.026 0.5993 9.602 Mn -18155.45 0.023 0.6194 0.013 Cl -22356.60 12.166 0.7139 Ca 591.97 -0.272 0.009 0.7603 -7287.41 6.085 0.001 Mg 0.9377 Al -1305.26 1.057 0.000 0.9426
- 826 highlighted in green. Compare Figure A8a.

828

Figure A1: Boxplots showing the distributions of physical stream water parameters (i.e. electrical conductivity, pH, redox potential and temperature) measured using a multiprobe sensor at station WU14. The box shows the median and the interquartile range and the whiskers extend to 1.5 times the interquartile range away from the box.

833

834

Figure A2: Boxplots showing the distributions of physical stream water parameters (i.e.
electrical conductivity, pH, redox potential and temperature) measured using a multiprobe sensor at station WU17. The box shows the median and the interquartile range and

the whiskers extend to 1.5 times the interquartile range away from the box.

839

840 Figure A3: Boxplots showing the distributions of all considered macro- and

841 micronutrients, dissolved aluminum and DOC determined from stream water samples

taken at discharge station WU17. The box shows the median and the interquartile range

and the whiskers extend to 1.5 times the interquartile range away from the box.

Figure A4: Nitrate prediction models fitted to the measured values from weekly samples at
WU17 using functions implemented in the LOADFLEX package. Subplots show predicted
concentrations for rectangular interpolation (no3_interp), linear regression model (no3_lm),
RLOADEST regression model (no3_reg2), and composite model (no3_comp).

850

845

Figure A5: Predicted nitrate concentrations at WU14 using the four prediction models and the high-resolution nitrate concentrations measured with the TriOS sensor for the years 2014 (upper panel) and 2020 (lower panel). Subplots show observed concentrations (green dots) and predicted concentrations (orange lines) for rectangular interpolation (no3_interp), linear regression model (no3_lm), RLOADEST regression model (no3_reg2), and composite model (no3_comp).

859

858

Figure A6: Nitrate concentrations from event and grab-based samples (black dots) at WU14
(above panel) and WU17 (below panel), respectively, and from four prediction models (red
lines) implemented in the LOADFLEX package. Subplots show predicted concentrations
for rectangular interpolation (DOC_interp), linear regression model (DOC_lm),
RLOADEST regression model (DOC_reg2), and composite model (DOC_comp).

Figure A7: Predicted hourly concentrations of all considered solutes at WU14 using thecomposite model.

Figure A8: Annual average concentrations of the solutes considered at a) WU14 and b) WU17, and linear trend lines. The red dots indicate the years before deforestation.

Year

877

876

Figure A9: Annual fluxes of the solutes considered at a) WU14 and b) WU17, and lineartrend lines. The red dots indicate the years before deforestation.