

# Biologically effective daily radiant exposure for erythema appearance, previtamin D<sub>3</sub> synthesis and clearing of psoriatic lesions derived from erythema-biometerserythema broadband meters at Belsk, Poland, for the period 1976-2023

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**Abstract.** A long-term series of exposures to solar ultraviolet (~~UV~~) radiation (UVR) is required to assess the risks and benefits of radiation on different human biological processes. However, homogenisation of the amount of biologically effective solar energy ~~reaching the Earth's surface over long periods (i.e. (i.e.~~ energy weighted according to the sensitivity of the selected biological process to solar radiation) reaching the Earth's surface over long periods is challenging due to changes in measurement methods and instruments. This paper presents the world's longest homogenised time series of biologically effective daily radiant exposures (DRE) from regular monitoring with different erythema ~~biometers (EB~~ broadband radiometers (EBRs) operated at the Central Geophysical Laboratory of the Institute of Geophysics, Polish Academy of Sciences (IG PAS), Belsk (20.79°E, 51.84°N) from 1 January 1976 to 31 December 2023. The following biological effects were considered: the ~~appearance of~~ erythema, cutaneous synthesis of previtamin D<sub>3</sub>, and clearing of psoriatic lesions. The data for the latter two biological effects are estimated based on the proposed method of using ~~EBEBR~~ measurements to calculate other non-erythema DRE. The following ~~broadband erythema radiometers EBRs~~ were used in the monitoring: Robertson-Berger (1975–1992), Solar Light model 501 (1993–1994 with #927, 1995–2013 with #2011) and Kipp-Zonen UV-AE-T #30616 from 5 August 2013 to the present. From 1976 to 2013, the homogenisation procedure consisted of comparing the measured erythema DRE and daily maximum of UV index (~~erythema irradiance at noon~~ UVI<sub>MAX</sub>) with the corresponding synthetic values from simulations using a radiation transfer model: for cloudless conditions. Between 2014 and 2023, the ~~raw EBR~~ data were compared with data from a collocated reference instrument, the Brewer Mark II #64 spectrometer. Such comparisons resulted in a set of multipliers that were applied to the ~~raw EBEER~~ measurements. Two different versions of the homogenisation method were applied (~~analysing modelled and observed values~~ for erythema DRE and ~~UV index with UVI<sub>MAX</sub> assuming~~ different ~~selection of criteria for~~ cloudless days), and three. Three regression models ~~were constructed for of~~ the erythema data ~~based on~~ common UVR indices (total column ozone, aerosol optical depth and global clear sky irradiance ~~clearness index~~ index) were used to reconstruct the UVR data from the beginning of the Belsk observations, allowing further validation of the homogenised UVR data. Linear trends calculated ~~from reevaluated and reconstructed time series (a total of seven time series were considered)~~ showed a statistically significant increase in erythema annual and summer (June to August) radiant exposures of about 6 % per decade over the period 1976–2005. Thereafter, no trend was observed. The same trend estimates were found for all biological effects considered.

The ~~raw and reevaluated~~ data are made freely available via the following repository:  
<https://doi.org/10.1594/PANGAEA.972139> (Krzyścin et al., 2024). An additional version of the ~~reevaluatedre-~~  
~~evaluated~~ data, together with the corresponding clear sky and proxy data used in the ~~UVUVR~~ data reconstruction,  
 is archived at:  
[https://doi.org/10.25171/InstGeoph\\_PAS\\_IGData\\_Biologically\\_Effective\\_Solar\\_Radiation\\_Belsk\\_1976\\_2023](https://doi.org/10.25171/InstGeoph_PAS_IGData_Biologically_Effective_Solar_Radiation_Belsk_1976_2023)  
 (Krzyścin, 2024).

**Keyword(s):** ~~biometererythema~~ broadband radiometer; biologically effective irradiance, homogenisation, radiant exposure

## 1 Introduction

Molina and Rowland (1974), winners of the 1995 Nobel Prize in Chemistry, argued that man-made chlorofluorocarbons (CFCs), which were widely used in industry in the 1970s, could ~~penetrate~~reach the stratospheric ozone layer where they ~~were~~would be destroyed by short-wave ultraviolet (~~UV~~)radiation; (UVR), releasing free chlorine atoms and causing stratospheric O<sub>3</sub> depletion in the catalytic reaction cycle. Solar radiation in the shortest part of its spectrum that reaches the Earth's surface (290–315 nm), known as UV-B, is strongly absorbed by stratospheric ozone. The discovery of the ozone hole over Antarctica (Chubachi, 1984; Farman et al., 1985) and the predicted decreasing trend in total column ozone (TCO<sub>3</sub>) in other regions have stimulated interest in establishing continuous monitoring of UV-B irradiance reaching the ground. In addition, there is growing evidence that such ~~UVUVR~~ trends can cause various adverse health effects, such as skin cancers (including the deadly melanoma), DNA damage, immunosuppression, oxidative stress and skin ageing (Neale et al., 2023). Solar UV-B radiation from space is attenuated as it passes through the atmosphere due to light scattering (by cloud particles, atmospheric gases and aerosols) and absorption (by O<sub>3</sub>, NO<sub>2</sub>, SO<sub>2</sub> and aerosols). The attenuation of light increases with its path length through the atmosphere (i.e. usually described by the air mass), so solar elevation and ground surface altitude are key parameters to consider in surface ~~UVUVR~~ modelling. Other factors forcing ~~UVUVR~~ variability at the surface that are often used as proxies for atmospheric UV-B attenuation are total column O<sub>3</sub> (TCO<sub>3</sub>) to account for ~~UVUVR~~ absorption by ozone, the ~~clearness~~clear sky index (CI) (i.e. a quotient of the all-sky global solar irradiance (~~GSIG~~) at the surface and the corresponding synthetic clear-sky value (G<sub>0</sub>) to account for combined cloud/aerosol scattering effects on ~~UVUVR~~), and aerosol optical depth (AOD) in the solar UV range (parameterising ~~UVUVR~~ attenuation by aerosols). TCO<sub>3</sub> and ~~GSIG~~ have been found to be the most effective for modelling surface UV-B radiation (Koepeke et al., 2006, den Outer et al., 2010).

In the early 1970s, the broadband Robertson-Berger (RB) meter was developed to measure the biologically effective (BE) ~~UV-radiation~~UVR that causes skin redness, also known as erythema (Berger, 1976). The spectral characteristics of RB resembled the erythema sensitivity of human skin. RB instruments began continuous monitoring of erythema irradiance in 1974 at eight sites in the United States (Scotto et al., 1988). During the 1970s, instruments were operated in other countries (Austria, Australia, Germany, Poland, Sweden, Switzerland) (WMO, 1977). At the beginning of this global network, RB meters were calibrated using a travelling standard meter provided by the Photobiology Center at Philadelphia University. After a few years, at some stations, including the Institute of Geophysics, Polish Academy of Sciences (IG PAS) station at Belsk (51.84°N, 20.78°E), this calibration method was replaced by comparisons with values modelled by the radiative transfer model. The Dave-Halpern model (Dave and Halpern, 1976) was used to estimate erythemally weighted irradiance for cloudless

sky conditions to calibrate the Belsk data (Słomka and Słomka, 1985). Serious drawbacks of RB measurements were their results in relative units (counts), temperature sensitivity, a lot of manual work in data preparation, sometimes rapid ageing, and difficulties in accurately converting counts into the so-called sunburn unit (the minimum erythral radiation exposure that causes redness of the skin). These problems were significantly reduced in a new version of the RB meter, a prototype of the current UV biometer, erythral broadband radiometer (EBR), developed in the late 1980s as a result of collaboration between IG PAS and the Institute of Medical Physics of the University of Innsbruck (Blumthaler et al., 1989; Słomka and Słomka, 1993). Further prototype work at Solar Light (SL) Co. in Philadelphia resulted in the production of a commercial SL Biometer ModModel 501A, which replaced the RB meter.

Other EBR versions ~~of broadband UV biometers for UV monitoring~~ were introduced in the 1990s, including those from Yankee Environmental Systems (Turner Falls, USA) and Kipp and Zonen (KZ) Co. (Delf, Netherlands). However, there was a need to standardise the calibration correction procedure for the broadband UVUVR meters as it became apparent that the calibration provided by the manufacturer could not be relied upon even for the same type of instrument (Leszczynski et al., 1998). A standard calibration method that takes into account the individual spectral characteristics of the instrument and the loss of sensitivity has been proposed (Hülsen and Gröbner 2007). However, uncertainties of ~7 % can still be expected for well-maintained ~~biometers~~EBRs (Gröbner et al., 2009). Long-term series of surface UVR from ground-based observations with a length of at least a few decades are rare. To the authors' knowledge, the longest UVR monitoring series began in Moscow in 1968 with a broadband (300–380 nm) instrument developed at the Moscow State University Meteorological Observatory (Chubarova et al., 2000). One of the world's longest measurements of solar UVR at the Earth's surface (and probably the longest taken by erythral broadband instruments) are from Belsk. Measurements began in 1975 and continuous monitoring started on 1 January 1976. From a global perspective, the first UVR results appeared at the World Ozone and Radiation Data Centre (WOUDC) in 1989, but continuous UVR time series over three decades are only available for a limited number of stations including: Uccle (Belgium), Edmonton, Resolute, Toronto, Churchill, Saturna Island (Canada), Tateno (Japan) and Syowa (Antarctica) (WOUDC, 2025). Database Network for the Detection of Atmospheric Composition Change (NDACC) include also stations with at least of three decades of UVR measurements such as Lauder (New Zealand), Mauna Loa (USA) and three Antarctica stations – Arrival Heights, Palmer Station and South-Pole (NDACC, 2025).

This article presents a retrospective evaluation of all UVUVR measurements (1976–2023) at Belsk made with different ~~broadband instruments~~EBRs including: RB (1976–1992), SL ~~biometer~~Biometer model 501 A (SL501 A) (two instruments were used #927 and #2011 for the period 1993–1994 and 1995–2013, respectively) and KZ UV-AE-T #30616 (KZ616) from 5 August 2013 to the present. The ~~reevaluation~~re-evaluation for the period 1976–2013 is based on a comparison of the measurements with the synthetic daily erythral ~~irradiance~~radiant exposure and UV index ~~(the midday value of erythral irradiance)~~at noon from a radiative model simulation for clear sky conditions using TCO<sub>3</sub> and AOD measured at Belsk as model input parameters. The quality of the KZ616 data (2013–2023) will be accessed through comparisons with clear-sky erythral irradiances simultaneously measured by the well-maintained Brewer spectrophotometer Mark II #64 (BS64). The details of the Brewer maintenance can be found in Czerwińska and Krzyścin (2024a). Erythral daily radiant exposures (DRE) for the entire period of the UVUVR measurements at Belsk will be transferred to the corresponding vitamin D<sub>3</sub> and antipsoriatic DRE using a method proposed by Czerwińska and Krzyścin (2024a). A comparison of these DRE with those from BS64

spectral measurements in the period 2014–2023 will indicate the accuracy of the proposed reconstruction method of past BE data based on a statistical approach using typical proxies ( $\text{TCO}_3$ , ~~GSI~~ and G) characterising atmospheric ~~UVUVR~~ attenuation. Finally, trend calculations in annual (January–December) and summer (June–August) radiant exposures (RE) for all biological effects considered and versions of the recalculated ~~UVUVR~~ data from 1976–2023 will be presented to confirm the robustness of the long-term changes in the BE radiation measured at Belsk.

## 2 Materials and Method

### 2.1 ~~UVUVR~~ monitoring

~~The recording~~Recording of ~~solar-erythema~~erythema irradiance with a standard RB meter (detector recorder #No. 40), ~~initiated in Belsk~~ started in May 1975, ~~was carried out at Belsk, but continuous monitoring began on 1 January 1976 and lasted~~ until 1994. From May 1993, in parallel with the RB measurements, the monitoring of erythema irradiance using the SL Biometer 501 A #927 was initiated in order to establish monthly transfer coefficients for converting the RB output in sunburn units (SU) into erythema units, i.e. the minimum erythema dose (MED) causing skin redness in typical Caucasian skin, which was entered into the SL Biometer 501 A measurements (Puchalski, 1995). It was assumed that  $1 \text{ MED} = 210 \text{ J}_{\text{eryt}} \text{ m}^{-2}$ , where  $\text{J}_{\text{eryt}}$  denotes spectral irradiance integrated over time and wavelengths (290–400 nm) after weighting by the erythema action spectrum ~~(CIE, 2019)~~. Simultaneous measurements continued until December 1994, and all erythema DRE measured with the RB meter before 1993 were multiplied by these transfer coefficients to obtain data comparable to those with the SL Biometer 501 A.

As the RB meter showed sensitivity to ambient temperature, a correction for temperature effect was applied to the raw daily RB values (Borkowski, 1998) using empirical formulas proposed by Koskela et al. (1994). In addition, the RB Belsk series was also found to be affected by a change in calibration method in 1985, as ~~the~~ Dave-Halpern model ~~(Dave and Halpern, 1976)~~ calculations for cloudless conditions replaced field comparisons with the travelling standard instrument. This resulted in a downward step change of 14 % in the ~~UVUVR~~ series (Borkowski, 2000). The ~~reevaluated~~re-evaluated time series of erythema DRE for the period 1976–1992 as made by Borkowski (2008) was archived and formed part of the ~~raw~~ Belsk’s erythema time series (1976–2023), which is further homogenised in this study.

Subsequent ~~UVUVR~~ measurements included SL501 A # 927 (1993–1994) and #2011 (1995–2013), which were only ~~pre-roughly~~ calibrated by the instrument manufacturer, ~~prior to shipment~~. In 2005, KZ616 was added to the IG PAS UV network and served as the reference instrument. It was not used for everyday ~~UVUVR~~ monitoring but only for occasional international calibration campaigns to provide a source for further calibrations with our SL biometers operating in Belsk and Hornsund (Spitzbergen). KZ616 started regular ~~UVUVR~~ monitoring on 5 August 2013, replacing the ~~previous raw~~ SL501 A #2011, as BS64 (normally measuring  $\text{TCO}_3$  and Umkehr ozone at Belsk since 1992) was established as the new ~~UVUVR~~ reference instrument for the IG PAS network, which has been in operation until now. The performance of KZ616 has ~~proved~~proven to be very stable and ~~it~~ is still involved in regular ~~UVUVR~~ monitoring.

In the following, we use the term “raw data” for the results of the EBR measurements in  $\text{W m}^{-2}$  and  $\text{J m}^{-2}$  that were previously archived in the internal databases of IG PAS before the release of the Krzyścin et al. (2024) and Krzyścin (2024) databases.

## 2.2 Ancillary data

Daily representatives of  $\text{TCO}_3$  at Belsk are taken from the IG PAS data portal (Krzyścin, 2024), which contains results of daily average  $\text{TCO}_3$  measurements throughout the day, prepared for ~~UV~~UVR modelling purposes. For example, the most reliable daily representative value of  $\text{TCO}_3$  (marked with flag no. 1) was calculated as an average of the most accurate measurements (the so-called direct sun measurements) made by the Dobson spectrophotometer between 9:00 and 13:00 UTC. The least accurate case of ground-based  $\text{TCO}_3$  observations (with flag no. 5) occurred under cloudy and low sun elevation conditions, i.e. before 9:00 and after 13:00 UTC. In this case, only the least reliable Dobson observations were available for calculating the daily  $\text{TCO}_3$  representative under overcast zenith and high air masses. In the rare cases when ground observations were not available, satellite data (flag 6 or 7 depending on the data source) and/or  $\text{TCO}_3$  reanalysis data (flag 8) were used.

~~The daily representative of  $\text{CI}$ ,  $\text{DCI}$ , which is a commonly further used measure of cloud attenuation of global (direct and diffuse) solar irradiance at ground level (Liu and Jordan, 1960). Daily values of  $\text{CI}$  are in regression models (Sect. 2.3.4), is calculated as the quotient of the all-sky ( $G$ ) and the corresponding synthetic clear-sky (daily integrals (sunrise to sunset) of  $G$  and  $G_0$ ) daily integral of global solar irradiance. Typically,  $G$  is derived from observations and  $G_0$  from a model simulation, depending on the amount of solar absorbers (mostly water vapour) and AOD. Global solar DRE the former values were obtained from routine monitoring of global solar irradiance by various pyranometers (since 1965) including the following instruments: Kipp & Zonen CM 6, Sonntag PRM-2, Kipp & Zonen CM 5, Kipp & Zonen CM 11, and Kipp & Zonen CM 21. The data were calibrated using the Polish national standard, which washad previously calibrated been adjusted to the world standard during inter-comparison campaigns at the World Radiation Centre in Davos, Switzerland. In addition, the Campbell-Stokes sunshine recorder provided the duration of sunshine per day to pre-select sunny days. All these data are archived in the IG PAS Data Portal (Krzyścin, 2024).~~

To ~~support~~validate the ~~quality of the corrected~~ UV observations at Belsk, the long-term variability of BE ~~radiance~~radiation was also obtained from the ~~UV~~UVR reconstruction models (Section 2.3) using proxies ( $\text{TCO}_3$  and  ~~$\text{CI}$~~  $\text{DCI}$ ) from the ground-based observations and reanalysis datasets. The European Centre for Medium-Range Weather Forecasts (ECMWF) v5 (ERA5) reanalysis provides, in addition to many other variables, intra-day  $\text{TCO}_3$  values, ~~global solar irradiance for clear sky  $G_0$~~ , and ~~all-sky conditions  $G$~~  for the period 1940–2024, which are freely available on the ERA5 (~~2024~~2025) website. Also included are data (from 1 January 1980 to the present) downloaded from the Modern-Era Retrospective Analysis for Research and Applications version 2 (MERRA-2) database (GMAO, ~~2024~~2025) using the Giovanni data search tool, which is freely available on the Giovanni (~~2024~~2025) website.

**Table 1. The Belsk's instruments and their working periods.**

Data	Instrument/data	Operation period	Reference
<u>Daily ERE and UV Index</u>	<u>Robertson Berger Meter</u>	<u>1976–1994</u>	<u>Krzyścin et al. (2024)</u>
	<u>SL Biometer 501 A # 927</u>	<u>1992–1994</u>	
	<u>SL Biometer 501 A # 2011</u>	<u>1995–2013</u>	
	<u>Kipp-Zonen UV-AE-T # 30616</u>	<u>2013–present</u>	
<u><math>\text{TCO}_3</math></u>	<u>Dobson Spectrophotometer # 84</u>	<u>1963–present</u>	<u>Krzyścin (2024)</u>
<u>SunDur</u>	<u>Campbell–Stokes sunshine recorder</u>	<u>1966–1968,</u> <u>1970–1973,</u> <u>1975–present</u>	<u>Krzyścin (2024)</u>
<u><math>G</math></u>	<u>Kipp CM 6</u>	<u>1965–1980</u>	<u>Krzyścin (2024)</u>
	<u>Sonntag PRM-2</u>	<u>1981–1987</u>	

	<a href="#">Kipp&amp;Zonen CM 5</a>	<a href="#">1988–1991</a>	
	<a href="#">Kipp&amp;Zonen CM 11</a>	<a href="#">1992–2010</a>	
	<a href="#">Kipp&amp;Zonen CM 21</a>	<a href="#">2010–present</a>	
<a href="#">AOD<sub>340nm</sub></a>	<a href="#">Sonntag pyrheliometers</a>	<a href="#">1976–2013</a>	<a href="#">Krzyścin (2024)</a>
	<a href="#">CIMEL CE 318-T</a>	<a href="#">2004–present</a>	
<a href="#">G and G<sub>0</sub></a>	<a href="#">ERA5 reanalysis</a>	<a href="#">1940–present</a>	<a href="#">ERA5 (2025)</a>
<a href="#">G<sub>0</sub></a>	<a href="#">MERRA-2 reanalysis</a>	<a href="#">1980–present</a>	<a href="#">GMAO (2025)</a>

Atmospheric aerosols can be significant drivers of surface ~~UV radiation~~UVR, especially under clear sky conditions (Krzyścin and Puchalski, 1998). The column properties of aerosols can be obtained from ground-based observations and used in the modelling of radiative transfer in the atmosphere. Aerosol properties are described by various characteristics (e.g. including AOD, single scattering albedo, asymmetry factor). In this article, we use Belsk's AOD at 340 nm (IG PAS Data Portal, Krzyścin (2024)), which is estimated from the Linke turbidity factor measurements with Sonntag pyrheliometers between 1976 and 2013 (Posyniak et al., 2016) and from the co-located solar photometer CIMEL CE 318-T (~~2014~~2004–2023) operating within the Aerosol Robotic Network (AERONET) (AERONET, ~~2024~~2025). Other aerosol properties are kept constant and equal to their typical values for the rural site. [Table 1 summarises the sources of the data used in this paper.](#)

## 2.3 ~~UV~~UVR models

### 2.3.1 Clear-sky model

~~Radiative transfer model simulations for clear sky conditions are used to quantify and correct biases in the output of the Belsk UVR radiometers. To speed up the calculations, the look-up tables were obtained using the Tropospheric Ultraviolet and Visible (TUV) Radiative Transfer Model (TUV, 2025).~~

Synthetic clear-sky values of BE (erythema ~~appearance~~, previtamin D<sub>3</sub> synthesis, clearing of psoriasis lesions) RE ~~and in day D,  $RE_{EFF,CS}(D)$  in  $J_{EFF} m^{-2}$ , and irradiance at noon in day D,  $RE_{EFF,CS}(D)$  in  $J_{eff} m^{-2}$ , and  $Ir_{EFF,CS}(t=noon)$  in  $W_{eff}W_{EFF} m^{-2}$ , respectively, are derived from look-up tables obtained from the Tropospheric Ultraviolet and Visible (TUV) radiation transfer model (TUV, 2024);~~calculated using the following formulas:

$$RE_{EFF,CS}RE_{EFF,CS}(D) = \int_{Sunrise(D)}^{Sunset(D)} Ir_{EFF,CS} \int_{Sunrise(D)}^{Sunset(D)} Ir_{EFF,CS}(t) dt$$

(1)

$$Ir_{EFF,CS}Ir_{EFF,CS}(t) = \int_{290\text{ nm}}^{400\text{ nm}} Ir_{CS}(\lambda, t) AS_{EFF}(\lambda) d\lambda$$

(2)



where  $I_r(\lambda, t) - I_{r,CS}(\lambda, t)$  is the spectral irradiance at time  $t$  for the and at wavelength  $\lambda$ , and  $AS_{EFF}(\lambda)$  denotes the action spectrum for specific biological effect EFF: EFF=ERYT for erythema (CIE 2019), EFF=VITD3 for photosynthesis of previtamin D<sub>3</sub> in human skin (CIE 2006), and EFF=PSOR for psoriasis clearing (Krzyścin et al., 2012). Figure 1 presents the action spectra used.

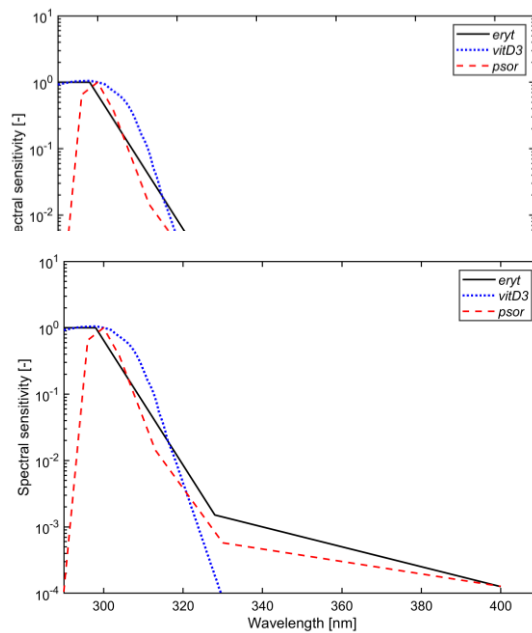
**Figure 1. Normalised action spectra for the specific biological effects: erythema appearance, (eryt), photosynthesis of previtamin D<sub>3</sub> in human skin, (vitD3), psoriasis clearing, (psor).**

Input to the clear-sky version of TUV model (daily representatives of TCO<sub>3</sub>, annual and monthly mean AOD at 340 nm for the period 1976–2013 and 2014–2023, respectively) and output ( $RE_{EFF, Clear-sky}(D)$  and  $I_{r, EFF, Clear-sky} RE_{EFF, CS}(D)$  and  $I_{r, EFF, CS}(t=noon)$ , where EFF={ERYT, VITD3, PSOR}), are archived in IG PAS Data Portal (Krzyścin, 2024).

### 2.3.2 Reevaluation of the UVUVR measurements

Model simulations of erythema DRE and noon UVI under cloudless sky provide a basis for the correction procedure of the past UVR data. A selection of clear-sky conditions throughout the entire day from the daily proxy values (relative sunshine duration and DCI), which were available for Belsk, is not straightforward as only the examination of the daily course of these measurements would allow to capture cloudless moments within the day. Thus, two different sets of correction coefficients are proposed, called CC1 and CC2.

The intraday UVUVR measurements at Belsk from 1976 to 2023 can be clearly divided into three periods:



1 January 1976–31 December 1992, 1 January 1993–4 August 2013, and 5 August 2013–31 December 2023, according to the different broadband instruments used for UVUVR monitoring, i.e. RB, SL501 A, and KZ616, respectively. For the first period, only the daily-erythema RE was DRE were archived, whereas for other periods the erythema irradiances at noon, the so-called daily maximum of UV index (UVI), were UVI<sub>MAX</sub> was also available (equal to the value of erythema irradiance at noon during a cloudless day). There were also periods when both instruments were operated simultaneously for calibration purposes: March 1992–December 1994 (RB

versus SL501 A), 5 August 2013–31 December 2014 (SL501 A #2011 versus KZ616), and 5 August 2013–31 December 2023 (KZ616 versus BS64).

The ~~calibrationcorrection~~ procedure before 5 August 2013 consisted of comparing the raw erythemal data with the corresponding synthetic values obtained from the radiative model simulations (described in Sections 2.3.1) for ~~the~~ days when ~~clear sky conditions can be assumed from the~~ ancillary data: ~~indicated that the sky was clear throughout the day.~~ The ~~locally weighted scatterplot smoother~~ Locally Weighted Scatterplot Smoothing (LOWESS; ~~proposed by~~ Cleveland, ~~(~~1979) was used to extract the smoothed pattern of the multipliers of the raw ~~UV~~ UVR data, i.e. the ~~calibrationcorrection~~ coefficients (CCs), from the daily ratios between synthetic and erythemal REs (for version CC1 of the ~~calibrationcorrection~~ or from the ratios between ~~UVIs~~ UVI<sub>MAX</sub> (version CC2) ~~taken for the days when clear sky conditions can be assumed at Belsk. Two sets of CCs were examined to determine the range of uncertainty in the CC estimates. In order to allow for greater variability in the CC values, different criteria for clear sky conditions were applied, and the smoothing procedure was applied to the long (1976–2013) and short (1993–2013) UV time series for the CC1 and CC2 versions, respectively. Accordingly, the~~ The following conditions were applied for the selection of clear sky ~~sets~~ data used in the two correction methods:

- CC1 – direct sun TCO<sub>3</sub> measurements occurred between 9:00–13:00 UTC (code 1 for the TCO<sub>3</sub> observation in IG PAS Data Portal, Krzyścin (2024) ) and the daily difference between the observed and the theoretical maximum sunshine duration ~~and theoretical one (for SZA < 85°)~~ is less than 0.5 hour as for higher SZA broadband UV30 minutes at solar zenith angles (SZAs) below 85°. This limit was chosen because broadband UVR measurements at larger SZAs are unreliable and the Campbell-Stokes instruments starts when direct sun irradiance exceeded 120 W m<sup>-2</sup>.
- CC2 – ~~For~~ for TCO<sub>3</sub>, the same condition was set as for CC1, and the ratio between the observed and theoretical sunshine hours (for SZA < 85°) is not less than 85 %. CC2 values have only been calculated for the period since 1 January 1993. ~~Prior to this date, we assumed that the calibration coefficients were equal to 1.0 according to the recalibration of the RB data in 2011 (Krzyścin et al., 2011).~~ Prior to this date, a re-evaluation of the RB data with a model mimicking the KZ radiometer measurements by Krzyścin et al. (2011) showed that the correction was not necessary, i.e. CC2=1. This choice is also confirmed here by the constant long-term patterns of CC1 in the period 1976-1992 (Fig. 6a), and only a small jump in the differences between CC1 and CC2 in 1992/1993 (Fig. 6b)

Different criteria for the selection of cloudless days would result in even greater differences between the two CC versions. In addition, the smoothing procedure was applied to the long (1976-2013) and short (1993-2013) UVR time series for the CC1 and CC2 versions, respectively. We would like to have two different sets of correction coefficients to find out how the long-term pattern of biologically effective radiation is sensitive to the corrections.

The CC1 and CC2 versions of the ~~reevaluated~~ re-evaluated Belsk ~~UV~~ UVR data are stored in the following free-access data archives: <https://doi.org/10.1594/PANGAEA.972139> (Krzyścin et al., 2024) and [https://doi.org/10.25171/InstGeoph\\_PAS\\_IGData\\_Biologically\\_Effective\\_Solar\\_Radiation\\_Belsk\\_1976\\_2023](https://doi.org/10.25171/InstGeoph_PAS_IGData_Biologically_Effective_Solar_Radiation_Belsk_1976_2023) (Krzyścin, 2024), respectively.

### 2.3.3 Reconstruction of BE radiation from the erythemal data



Broad-band instruments for measurement of the erythral irradiance can also estimate non-erythral irradiance by multiplying the erythral irradiance by the so-called conversion factors ( $CF_{EFF}$ ) derived from spectral ~~UV~~UVR measurements and/or radiative transfer simulations (Schmalwieser et al., 2022; Czerwińska and Krzyścin, 2024a):

$$Ir_{EFF}(t) = CF_{EFF}(TCO_3, SZA) \times Ir_{ERYT}(t), \quad (3)$$

where  $SZA$  denotes the solar zenith angle at time  $t$ . Following this concept, the daily radiant exposure for previtamin D<sub>3</sub> synthesis and psoriasis clearance ~~on the current  $D$  in year ( $YR$ ), month ( $MM$ ), and day, of month ( $DD$ )~~  $RE_{VITD3}(D)(YR, MM, DD)$  and  $RE_{PSOR}(D)(YR, MM, DD)$ , respectively, were estimated using the daily conversion factor,  $CF_{EFF}^*$ , applied to the ~~reevaluated~~ erythral DRE: ( $RE_{ERYT}(YR, MM, DD)$ ):

$$RE_{EFF}(D)(YR, MM, DD) = CF_{EFF}^*(TCO_3, D^*) \times RE_{ERYT}(D)(TCO_3, JD) \times RE_{ERYT}(YR, MM, DD), \quad EFF = \{VITD3, PSOR\}, \quad (4)$$

where  $CF_{EFF}^*$  depends on  $TCO_3$  and  $D^*/JD$  (Julian day ~~of the year (i.e. between 1 and 365/366) number~~ corresponding to the current  $D$ -day:  $\{YR, MM, DD\}$ ).  $CF_{EFF}^*$  and  $CF_{EFF}$  values were obtained from the radiative transfer model simulations. The time series (1976–2023) of ~~these values and the conversion factors~~,  $RE_{EFF}(D)$  and  $(YR, MM, DD)$ , and the corresponding noon value of the biologically effective irradiance,  $Ir_{EFF}(t = \text{noon})$  ~~from Eq. (3–4) ( $t = \text{noon}$ )~~, have been archived in the IG PAS Data Portal (Krzyścin, 2024).

#### 2.3.4 Regression models

~~Various~~The CCs described in section 2.3.2 were obtained for cloudless conditions and applied to all-sky conditions, where the contribution of the diffuse part of the radiation increases with cloud cover and dominates under overcast conditions. It cannot be excluded that the instruments used to monitor UVR at Belsk have their own specific characteristics for recording diffuse radiation, and that  $CF_{EFF}^*$  and  $CF_{EFF}$  in Eqs. (3–4) should also depend on the combined characteristics of clouds and instruments. To test whether this is the case, we investigated how different regression models ~~built from~~, which were trained using the ~~UV~~UVR data collected ~~in the~~between 2014 and 2023 (for this period ~~2014–2023 allowed for extended~~, the quality of the broadband radiometer was confirmed by the Brewer Mark II observations), reproduce the daily doses of erythral ~~RE analysis for radiation~~ throughout the entire 1976–2023 monitoring period ~~to provide a quality measure of the reevaluated UV~~. The first model (Mod1) is based on clear-sky spectra determined with the RT model discussed in Section 2.3.1 and a cloud modification factor (CMF) derived from DCI data. The second and third models (Mod2 and Mod3) are based on  $TCO_3$  and DCI data evaluated on a monthly basis.  $TCO_3$  and DCI were either taken from observations at Belsk (Mod2) or ERA5 reanalysis (Mod3).

According to a ~~frequently~~widely used ~~UV~~UVR modelling concept (e.g. Rieder et al., 2008; ~~den~~ Outer et al., 2010; Čížková et al., 2018; Czerwińska and Krzyścin, 2024b) ~~that~~ the erythral DRE on the current day  $D$ :  $\{YR, MM, DD\}$ ,  $RE_{ERYT}(D)(YR, MM, DD)$ , is the product of ~~the so-called cloud modification factor (CMF), which is an~~CMF (empirical function of ~~CI~~DCI parameterising UVR attenuation by clouds) for that day, and the synthetic clear-sky value,  $RE_{ERYT,CS}(D)$  (Section. 2.3.1):

$$RE_{ERYT}(D) = CMF(CI(D))(YR, MM, DD) = CMF(DCI(YR, MM, DD)) \times RE_{ERYT,CS}(D), \quad (5)$$

$CMF(D)$  is ~~parameterised~~calculated here as a power function with the regression coefficients,  $\alpha$  and  $\beta$ , depending on SZA at noon,  $SZA_N$ , for the current day  $D = \{YR, MM, DD\}$ :

$$CMF(D)(DCI(YR, MM, DD)) = \alpha[DCI(YR, MM, DD)]^\beta, \quad (6)$$

where estimates for the regression coefficients,  $\alpha$  and  $\beta$ , were obtained from the 2014–2023 data when the KZ616 measurements were well-fitted matched to the BS64 data (Section 3.1). In ~~CI calculations~~ ( $CI = GG_0 DCI$  ( $DCI = DG DG_0^{-1}$ ), calculation, the daily integral of global solar ~~DRE,  $G$ , comes~~irradiance,  $DG$ , is from observations at Belsk or ERA5, and its clear-sky equivalent,  $G_0 DG_0$ , from ERA5 (before 1980), and thereafter the mean of ERA5 and MERRA-2 values.

The standard least-squares subroutine (Matlab function –  $fitlm(x,y)$ ) provided the estimates for three arbitrarily selected SZA ranges (Table 42). These regression coefficients were used for the reconstruction of the  $RE_{ERYT}(D)$  time series for the entire period of ~~UV~~UVR measurements (1 January 1976 up to 31 December 2023). This model will be referred to as Mod1 in the following text.

**Table 42.** Estimates of the regression coefficients,  $\alpha$  and  $\beta$ , describing the attenuation by the cloud of erythema DRE by the empirical model, Mod 1, defined by Eqs. (5–6), for the three SZA ranges of at noon SZA according to Eq. (6(SZAn)).

Regression Coefficients					
$\alpha$	$\beta$	$\alpha$	$\beta$	$\alpha$	$\beta$
SZAN < 45°		SZAN ≥ 45° and < 60°		SZAN ≥ 60°	
0.954	0.844	0.918	0.750	0.960	0.697

The next two regression models were ~~built~~trained using the monthly averages of erythema DRE,  $RE_{ERYT}(YR, M)(YR, MM)$ , for month  $M$  in year  $YR$  (from 2014 up to 2023) averaging all available daily  $RE_{ERYT}(D)(YR, MM, DD)$  values in  $M$  month for  $YR$  year. The corresponding long-term (2014–2023) monthly means for  $M$  month,  $RE_{ERYT}^*(M)(MM)$ , is from the averages of all data for this calendar month. The idea of these models is to explain relative changes in the erythema monthly RE, i.e.,  $\Delta ER(YR, M) = 100\% (RE_{ERYT}(YR, M)(YR, MM) - RE_{ERYT}^*(M)(MM)) / RE_{ERYT}^*(M)(MM)$  with the corresponding relative changes in the UV explaining variables  $X$ , i.e., that affect UVR,  $\Delta X(YR, M) = (YR, MM) = 100\% (X(YR, M)(YR, MM) - X^*(M)(MM)) / X^*(M)(MM)$ , where  $X = \{G, X(YR, MM)$  is the monthly mean of  $DG$  or  $TCO_3\}$  in year  $YR$  and month  $MM$  and  $X^*(MM)$  is the long-term monthly means for month  $MM$

$$\Delta ER_K(YR, M)(YR, MM) = a_K(M) \Delta G(YR, M)(M) \Delta DG(YR, MM) + b_K(M)(MM) \Delta TCO_3(YR, M)(YR, MM) + c_K, \quad (7)$$

where  $K = \text{OBS}$  (for Mod2) and ERA5 (Mod3) are for the regression using the explaining variables from the measurements at Belsk and ERA5 reanalysis, respectively. Finally, the modelled  $RE_{ERYT, K}(YR, M)$  value is equal to:

$$RE_{ERYT, K}(YR, M) = RE_{ERYT, K}^*(M) \left( 1 + \frac{a_K(M) \Delta G(YR, M) + b_K(M) \Delta TCO_3(YR, M) + c_K}{100} \right), \quad (8)$$

$$RE_{ERYT, K}(YR, MM) = RE_{ERYT, K}^*(MM) \left( 1 + \frac{a_K(MM) \Delta DG(YR, MM) + b_K(MM) \Delta TCO_3(YR, MM) + c_K}{100} \right), \quad (8)$$

Models defined by Eq. (8) were used to compare fluctuations in ~~UV~~UVR data in periods with RB and SL501 A measurements relative to the long-term monthly means in these periods,  $RE_{ERYT}^*(M) RE_{ERYT, K}^*(MM)$ , that were

approximated using the long-term averages of the measured  $RE_{ERYT}(YR, MM, DD)$  values for the period 1976–1992 and 1993–2013, respectively. The regression coefficients,  $a_K$ ,  $b_K$ , and  $c_K$ , which were calculated using the standard least-squares linear fit to the most reliable (2014–2023) data (Table 23), were applied to construct monthly time series for the entire measurement period (1976–2023). The model for K=OBS and ERA5 in Eq. (8) is denoted further in the text as Mod2 and Mod3, respectively.

**Table 23.** Coefficients of the multilinear regressions derived for each calendar month based on the explaining variables from the measurements at Belsk (Mod2) and ERA5 reanalysis (Mod3) data for the period 2014–2023.

Month:	Mod 2			Mod 3		
	$a_{OBS}$	$b_{OBS}$	$c_{OBS}$	$a_{ERA5}$	$b_{ERA5}$	$c_{ERA5}$
January	0.84	−0.77	−5.69	1.34	−1.22	−8.38
February	0.81	−1.12	−0.12	0.95	−1.40	−0.05
March	0.59	−0.93	−0.65	0.84	−0.98	−0.77
April	0.90	−0.85	−1.94	1.26	−1.22	−3.77
May	0.86	−2.00	1.14	0.86	−1.97	0.64
June	1.08	−0.87	−0.05	1.14	−0.83	0.11
July	0.69	−0.84	0.00	0.40	−0.99	−0.00
August	0.82	−1.46	−1.99	0.63	−2.05	−1.40
September	0.86	−0.79	−0.00	0.94	−0.97	−0.00
October	0.80	−1.12	−0.49	0.86	−0.45	−0.52
November	0.58	−1.15	−1.02	0.66	−0.73	−0.97
December	0.73	−0.23	2.11	1.28	2.61	0.77

## 2.4 Statistical methods

Several standard statistical characteristics, which are calculated from the relative differences,  $z_i$ , between the observed,  $x_i$ , and model value,  $y_i$ , values expressed in percentage of the observed value, are used to determine the level of agreement between two time series. These are as follows: mean relative error (MREdeviation (MRD)), mean absolute error (MAEdeviation (MAD)), standard error (SEdeviation (SD)), root mean square error (RMSEdeviation (RMSD)), and Pearson's correlation coefficient (R):

$$z_i = 100\% \frac{y_i - x_i}{x_i}, \quad i = 1, \dots, N, \quad (9)$$

$$MREMRD = \frac{1}{N} \sum_{i=1}^N z_i, \quad (10)$$

$$MAEMAD = \frac{1}{N} \sum_{i=1}^N |z_i|, \quad (11)$$

$$SD = \left( \frac{1}{N} \sum_{i=1}^N (z_i - MRE)^2 \right)^{\frac{1}{2}} = \left( \frac{1}{N} \sum_{i=1}^N (z_i - MRD)^2 \right)^{\frac{1}{2}}, \quad (12)$$

$$RMSERMSD = \left( \frac{1}{N} \sum_{i=1}^N z_i^2 \right)^{\frac{1}{2}}, \quad (13)$$

$$R = \frac{\sum_{i=1}^N (x_i - \langle x \rangle)(y_i - \langle y \rangle)}{\left( \sum_{i=1}^N (x_i - \langle x \rangle)^2 \right)^{\frac{1}{2}} \left( \sum_{i=1}^N (y_i - \langle y \rangle)^2 \right)^{\frac{1}{2}}} = \frac{\sum_{i=1}^N (y_i - \langle y \rangle)(x_i - \langle x \rangle)}{\left( \sum_{i=1}^N (y_i - \langle y \rangle)^2 \right)^{\frac{1}{2}} \left( \sum_{i=1}^N (x_i - \langle x \rangle)^2 \right)^{\frac{1}{2}}}, \quad \langle x \rangle = \frac{1}{N} \sum_{i=1}^N x_i, \quad \langle y \rangle = \frac{1}{N} \sum_{i=1}^N y_i, \quad (14)$$

Standard least-squares linear regression is applied to find the long-term tendency in the data. According to Weatherhead et al. (1998), the standard error of the linear trend estimate,  $SE_{LS}$ , by standard least-squares approach should be multiplied by the factor  $F = \sqrt{(1+R_{k+1})/(1-R_{k+1})} \sqrt{(1+R_{k+1})/(1-R_{k+1})}$  to obtain the standard error corrected for the autocorrelation (with a time lag of 1) in the trend residuals,  $SE_{LS,COR}$ , if the trend residuals are positively correlated with the autocorrelation coefficient equal to  $R_{k+1}$ . ~~(for  $R_{k+1} < 0$ ,  $F=1$ ).~~ F is set to 1 if the autocorrelation coefficient in the residual time series is negative.

$$SE_{LS,COR} = F \times SE_{LS}, \quad (15)$$

Further in the text (Section 3.3), the slopes of the regression line will be calculated by Matlab function  $\text{fitlm}(x,y)$ , and the corrected standard error of the slope,  $SE_{LS,COR}$  for cases with  $R_{k+1} > 0$ , will be enlarged by the factor proposed by Weatherhead et al. (1998) (see Eq. (15)).

### 3 Results

#### 3.1 The ~~reevaluation~~ re-evaluation of the ~~UV~~ UVR measurements since 5 August 2013

On 5 August 2013, the KZ616 replaced the ~~previous~~ raw SL501 A #2011, which had been routinely used for UVR monitoring since 1995, as its performance had deteriorated (Fig. 2). Following this change, a new calibration ~~correction~~ procedure for the Belsk's biometer data UVR meter was introduced for early detection of instrument failure. Each month its output (erythemal irradiance) was compared with the corresponding output of the collocated BS64. An example of such a monthly comparison (for June 2023) ~~and time series of the monthly means of the ratio between BS064 and KZ616 erythemal DRE are shown in Fig. 3a and Fig. 3b, respectively.~~ is shown in the scatter plot between the BS064 and KZ616 erythemal irradiances measured under clear-sky conditions (Fig.3a). In addition, Fig. 3b shows the monthly ratios between these clear sky erythemal irradiances for the entire BS and KZ comparison period (2014-2023).

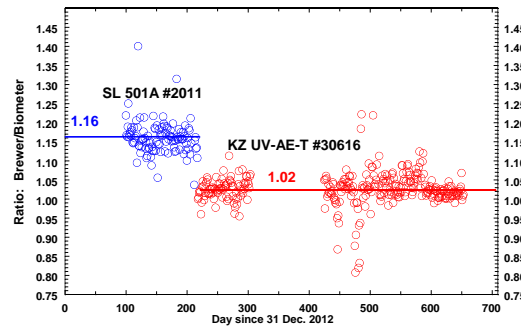


Figure 2. The ratio between the erythral DRE from the biometerservythermal radiometers (SL501 A #2011 before 5 August 2013 and KZ616 afterwards) and the Brewer Mark II spectrophotometer for the 2013–2014 period. The horizontal lines denote the mean value of the ratio.

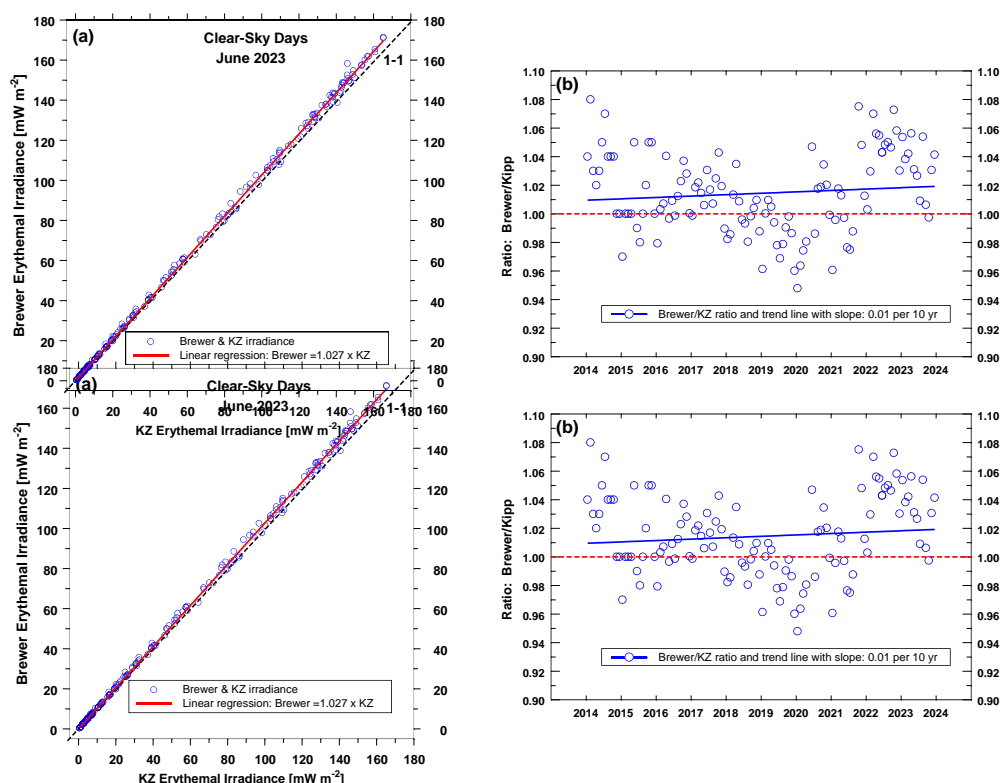
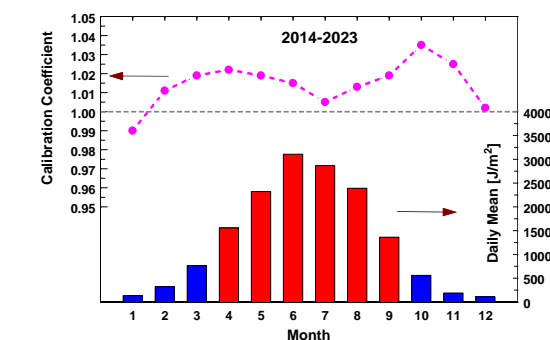
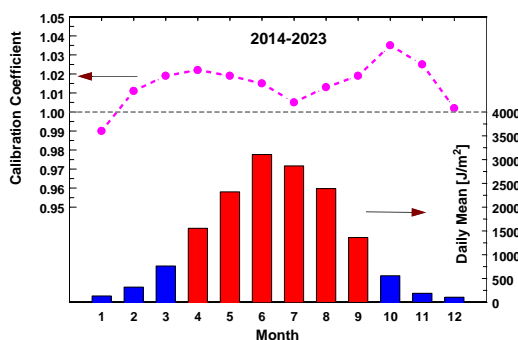


Figure 3. Comparison of the BS64 and KZ616 erythral data for the period 2014–2023: (a) the ratio between the erythral irradiances measured by the BS64 versus corresponding output of KZ 616 in June 2023 for clear-sky days,

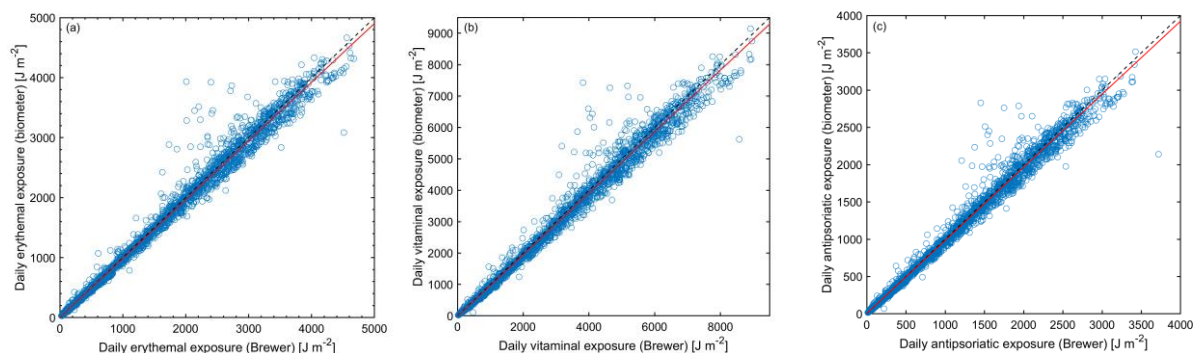


(b) time series of the monthly BS64/KZ616 ratios for the 2014–2023 period. The dashed line in Fig.3a shows the ideal 1:1 line.

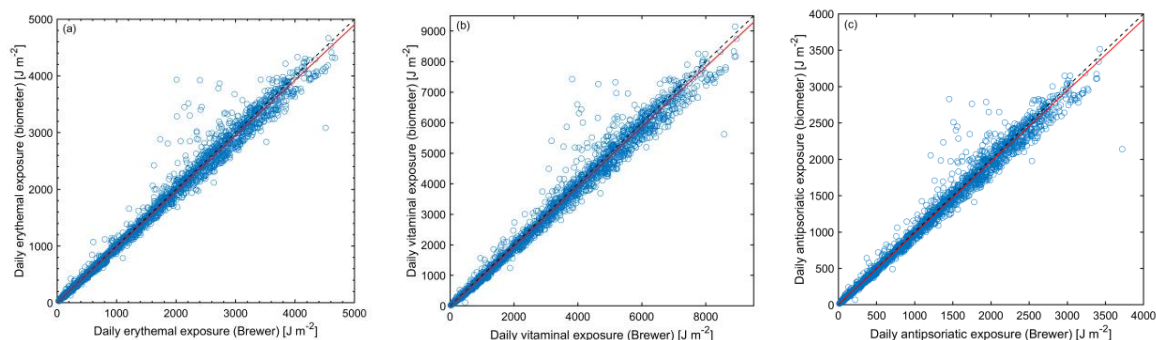


**Figure 4. Seasonal pattern of the ~~calibration~~correction coefficient (~~CC-ver-1~~CC1) and daily erythema RE for the period 2014–2023. Red bars denote months contributing mostly to the annual RE.**

The long-term (2014–2023) means of the monthly CC1 and erythema DRE for each calendar month are shown in the upper and lower graphs of Fig. 4. The ~~CC~~CC1 values are in the range of 1.00 to 1.02 during the period (April–September) when the intensity of solar ~~UV-radiation~~UVR is usually high and the fine weather often allows



prolonged outdoor activity. Given this and the insignificant trend in the time series of the monthly BS64/KZ ratios (Fig. 3b), it was decided to keep the original KZ616 data without additional adjustments. This assumption is also supported by the BS64/KZ616 comparisons ~~for all BE data considered for the period 2014–2023~~, as shown by the ~~linear regressions close to the 1–1 perfect agreement line~~ in the ~~three~~-scatter plots ~~(of Fig.-5)~~, which indicate that ~~the Brewer and EBR data cluster about the ideal 1:1 line~~. For the daily vitamin D<sub>3</sub> and antipsoriatic RE, the values were reconstructed from the daily erythema RE using the transfer coefficients defined by Eq. (4) (the values are archived in the IG PAS Data Portal, Krzyścin (2024), but the corresponding Brewer values were calculated from the real measured spectra weighted with the action spectra shown in Fig.1).



**Figure 5. Scatter plots (KZ616 versus BS64) for biologically effective DRE in the period 2014–2023: (a) erythema appearance, (b) previtamin D<sub>3</sub> synthesis, and (c) psoriasis clearance.**

Table A1 shows the values of the descriptive statistics for the period 2014–2023 according to the different ranges of SZA values at noon ( $SZA_N$ ), which confirm the good agreement between the DRE for all considered biological effects from the well-calibrated BS64 and KZ616 measurements used in routine UVUVR monitoring. For example, regardless of the biological effect, MREMRD and RMSEMRD are  $\sim -1\%$  and  $\sim 9\%$  for  $SZA_N < 45^\circ$ , which occurs from 8 April to 5 September at Belsk, i.e. during the period with the highest UVUVR intensity of the year. For  $SZA_N \geq 60^\circ$  (from 15 October up to 27 February of next year), MREMRD and RMSEMRD are only slightly larger ( $\sim -2\%$  and  $\sim 10\%$ , respectively) for the erythema and antipsoriatic exposures. These values are higher ( $\sim -13\%$  and  $\sim 18\%$ ) for the previtamin D<sub>3</sub> exposures, raising questions about the usefulness of the



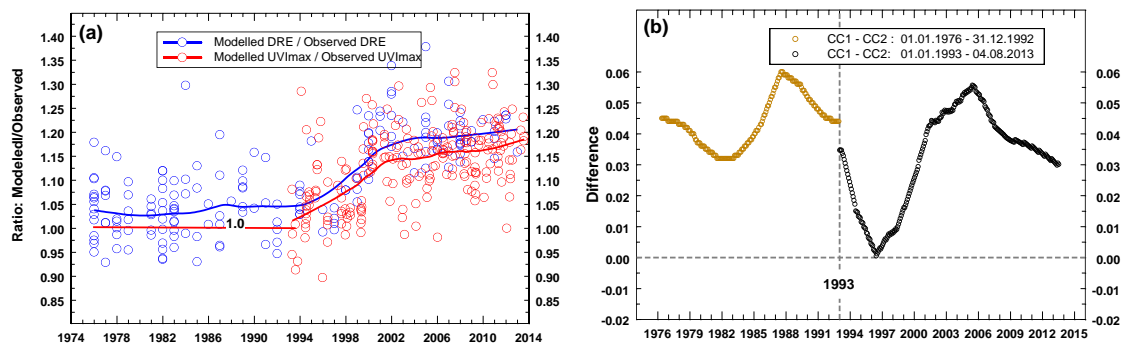
erythema ~~biometers~~radiometers for measuring vitamin D<sub>3</sub> exposure- when  $SZA_N \geq 60^\circ$ . However, vitamin D<sub>3</sub> synthesis in the skin ceases during this period.

### 3.2 The ~~reevaluation~~re-evaluation of the ~~UV~~UVR measurements before 5 August 2013

#### 3.2.1 ~~Calibration~~Correction coefficients from the clear-sky model simulations

Analyses of intraday ~~UV~~UVR measurements in Belsk from 1 January 1976 to 4 August 2013 have to be divided into two parts, i.e. 1 January 1976–31 December 1992, and 1 January 1993–4 August 2013, due to the different broadband instruments used for ~~UV~~UVR monitoring. In the first period, daily erythematous exposures were archived on the basis of manual summation of RB counts per day. For the latter period, 1-min erythematous irradiances were automatically recorded by a logger using SL501 A biometers and utilized in the calculation of ~~UVI~~UVI<sub>MAX</sub> and daily erythematous RE. Two methods of data ~~calibration for the period 1976–2013~~are correction were proposed (Sect. 2.3.2) using clear-sky data: modelled and measured daily erythematous RE ~~and UVI (for the period 1976–2013) and UVI~~MAX (1993–2013) for the correction method denoted CC1 and CC2, respectively. Figure 6a shows the time series of CC1 and CC2 values together with their smoothed values by the ~~LOeally Weighted Scatterplot Smoothing~~ (LOWESS) smoother, ~~Cleveland (1979)~~, which were used as multipliers of the raw ~~UV~~UVR data before 5 August 2013. The ~~differences~~difference between CC1 and CC2 are shown in Fig. 6b.

In the ~~former~~1976–1992 period, UVI values were not archived. This means that CC2 values cannot be directly calculated. However, CC2 values equal to 1 could be assumed as the output of the RB instrument was previously adjusted to that by SL501 A #927 using their simultaneous measurements for the period 1992–1994 (Puchalski et al., 1995). Such an assumption can also be supported here by a small jump (~1-%) in the ~~differences~~difference between the CC1 and CC2 values in January 1993 (Fig. 6b). This jump is really small taking into account that the 1993 adjustment of RB meter was inferred from field comparisons between RB and SL501 A #927 but here this



is calculated from smoothing ratios between modelled and observed UVI at noon for clear-sky days. Moreover, in the period 1976–~~1993~~1992, an oscillation with 0.015 amplitude is seen around the constant level of CC1=1.045 which justifies the assumption of an almost constant CC2 pattern before 1993. Using two sets of the ~~reevaluated~~re-evaluated 1976–2013 data will allow us to discuss the robustness of trend calculations for the entire 1976–2023 period of the ~~UV~~UVR measurements at Belsk (Sect. 3.3).

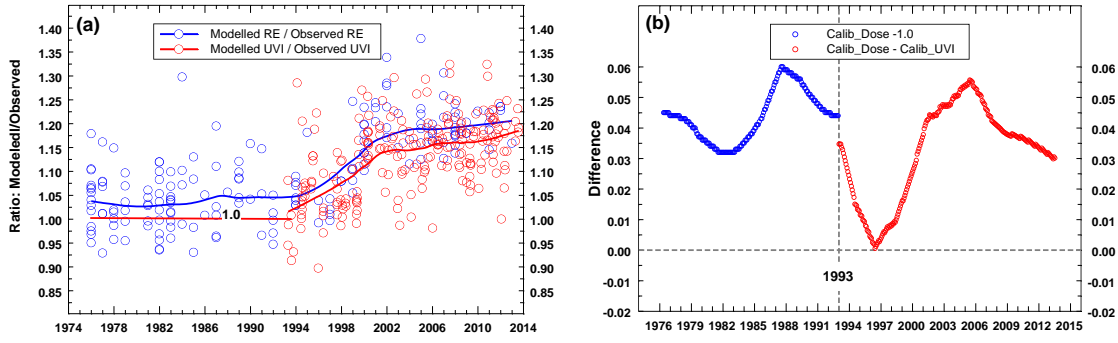


Figure 6. (a) TUV model-observation ratios for erythral DRE and noon UVI obtained for clear-sky days. The solid curves represent smoothed values of the ratios to be used as the calibration correction coefficients, i.e., the multipliers applied to the raw measurements. The multipliers were set equal to 1 for the 1976–1993 calibration 1992 correction based on UV ratios of noon UVIs, (b) differences difference between the monthly means of the calibration correction coefficients shown in Fig.6a.

### 3.2.2 Statistical Performance of the regression models

Erythral DRE for the period 1 January 1976 – 4 August 2013 were reconstructed with Mod1 defined by Eq.(Eqs. (5)–(6)). The model's constants came from the model training using the original KZ data and the explaining variables ( $\text{TCO}_3$  and  $\text{CHDCI}$ ) from 5 August 2013 – 31 December 2023 period. The reconstructed values were compared with two sets of the reevaluated re-evaluated data obtained before 5 August 2013 after multiplying raw daily erythral RE with CC1 and CC2, respectively.

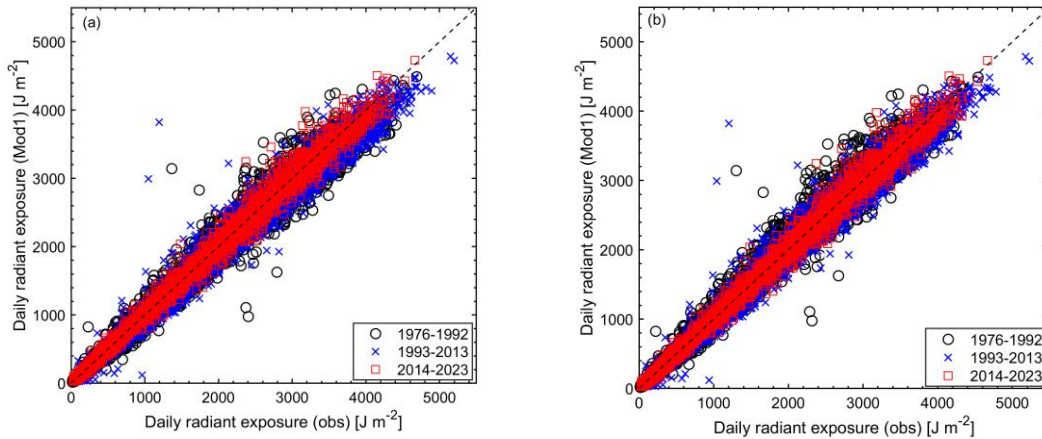
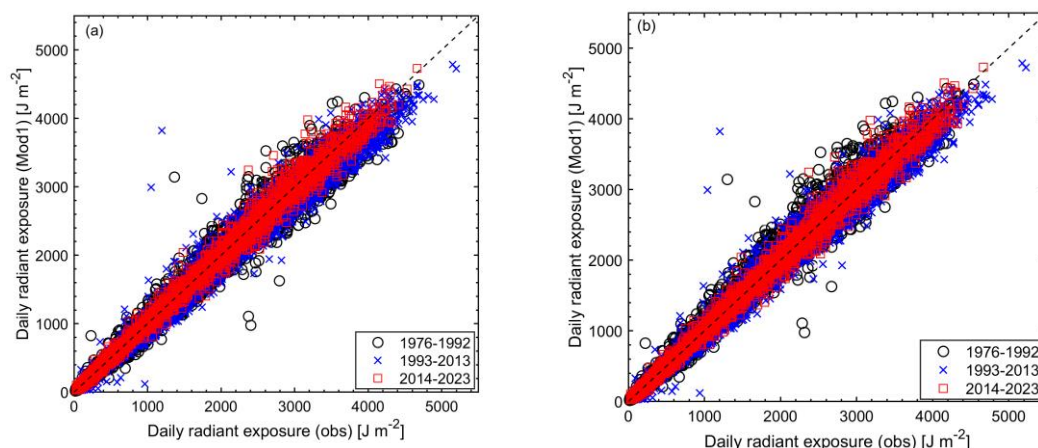


Figure 7. Scatter plot of the modelled (Mod1) erythral DRE versus the re-evaluated observed values for the 1976–1992, 1993–2013, and 2014–2023 period, respectively: (a) CC1 version of the correction coefficients for the period 1 January 1976–4 August 2013, (b) corresponding CC2 version of the correction coefficients. KZ616 measurements were taken without corrections.

Figure 7 shows the scatter plot of the reconstructed (Mod1) versus reevaluated re-evaluated erythral DRE with CC1 (Fig.7a) and CC2 (Fig.7b) multipliers of the raw data for the three periods corresponding to the RB, SL501A, and KZ616 measurements, respectively. The points in Figure 7 cluster around a line of perfect 1-1 agreement with only a few outliers. It seems that there is only a small difference between the reevaluated re-evaluated daily erythral RE and the corresponding output of Mod1 with the CC1 and CC2 multipliers. This is also supported by similar values of the descriptive statistics for the periods 1976–1992 and 1993–2013 (Table 34). It is worth mentioning that the performance of Mod1 resembles in the period 2014–2023 is similar to that of the Brewer

spectrophotometer from the comparison that was found when compared with the original KZ616 data (see almost note the same close values of the descriptive statistics for the full-year-round data in column “CC=1” of Table 34 and Table A1 for the “All SZA<sub>N</sub>-SZA<sub>N</sub>” cell, and column “Eryt”, for example, RMSE/RMSD values are equal to 10.5 % and 8.9 %, respectively).



**Figure 7.** Scatter plot of the modelled (Mod1) erythema DRE versus the reevaluated observed values for the 1976–1992, 1993–2013, and 2014–2023 period, respectively: (a) CC1 version of the calibration coefficients for the period 1 January 1976–4 August 2013, (b) corresponding CC2 version of the calibration coefficients. KZ616 measurements were taken without corrections.

**Table 34.** The descriptive statistics (MRE, MAE, RMSEMRD, MAD, RMSD, and SD, as defined in Sect. 2.4) calculated from the relative daily differences, 100% ((reevaluated Mod1 value – re-evaluated measurement – Mod1 value)/(reevaluated measurement)), for the periods 1976–1992, 1993–2013 and 2014–2023. The correlation coefficient R was obtained from the reevaluated re-evaluated measurements and modelled values. Two versions of the reevaluated re-evaluated datasets were considered, using CC1 and CC2 multipliers on the raw measurements. Both datasets include raw KZ616 data as there was no need to recalculate these data. The results are shown for annual (January–December) and summer (June–August) data.

Statistics	Year-Round (January–...–December)					June–July–August				
	Multipliers of the raw data									
	1976–1992		1993–2013		2014–2023	1976–1992		1993–2013		2014–2023
	CC1	CC2	CC1	CC2	CC=1	CC1	CC2	CC1	CC2	CC=1
MREMR	–2.7	–1.6	–1.9	–0.3	–1.4	–0.8	–3.5	–2.6	–1.0	–0.9
D					6.8					
MAEMA	9.8	9.6	9.7	9.4	10.5	7.8	8.1	7.0	6.4	5.2
D					10.0					
RMSEMR	13.7	14.1	14.5	14.6	1.00	10.8	11.7	10.1	9.7	6.9
SD					10.4					
R	1.00	0.99	1.00	0.99		0.96	0.96	0.97	0.98	0.98
SD	13.4	14.0	14.3	14.6		10.9	11.2	9.8	9.7	6.8

Erythema DRE by Mod1 can be obtained for days when the explanatory variables, TCO<sub>3</sub> and  $\text{C}_{\text{DCL}}$ , are available from the collocated measurements at Belsk by the Dobson radiometer and pyranometer, respectively. It is therefore possible to fill gaps in the measured data and obtain a complete (1976–2023) series of erythema DRE to be used in calculations of erythema annual and summer (June–July–August) RE. These data REs can also be calculated using the erythema monthly RE based on Mod2 and Mod3. All these series are analysed in section Section 3.3 for trend calculations to assess the level of uncertainty in the long-term variability of the Belsk UVUVR data.

Table 45 shows the values of the descriptive statistics for the three models used (Mod1, Mod2 and Mod3) and two versions of the ~~reevaluated~~re-evaluated data (using CC1 and CC2 multipliers on the raw data) based on the annual and summer RE. The differences between descriptive statistics (~~MRE, MAE, RMSEMRD, MAD, RMSD~~, SD) in CC1 and CC2 columns are within a few percentage points for ~~MREMRD~~ and about 1–1.5 percentage points for other statistics, indicating that the ~~two-independent-calibration~~different correction methods give fairly similar results. The performance of Mod2 and Mod3 is in most cases slightly better than that of Mod1 (Table 45) because these models add fluctuations to the mean values for the periods 1976–1992, 1993–2013 and 2014–2023 calculated from the ~~reevaluated~~re-evaluated measurements of RB, SL501A (#919 and #2011 for the periods 1993–1994 and 1995–2013 respectively) and the original KZ616 measurements.

All models considered were designed to test whether changes in the primary ~~UVUVR~~ drivers, ozone and clouds, explain year-to-year ~~UVUVR~~ variability. The performance of Mod3 is surprisingly similar to that obtained from Mod2 despite the use of ~~UVUVR~~ proxies (~~TCO<sub>3</sub> and DGI~~) from the ERA5 reanalysis. ~~This confirms the possibility of using explanatory variables from these reanalyses to fill gaps in the proxy data.~~

The lowest correlation coefficients between the ~~reevaluated~~re-evaluated measurements and modelled values were found in the period 1993–2013 for the measurement-model pairs with the same version of the CC multipliers (CC1 or CC2). This is particularly pronounced for the summer data (see e.g. Mod3 values of 0.50 and 0.43 for CC1 and CC2 pairs, respectively), suggesting a poorer agreement between measurements and model in the period 1993–2013. This was found for all models. However, other descriptive statistics (~~MRE, MAE, RMSEMRD, MAD, RMSD~~ and SD) differed only slightly, i.e. less than 1.5 percentage points, ~~when values in CC1 and CC2 columns were compared.~~

**Table 45.** Same as Table 34, but the descriptive statistics are calculated using time series of erythemal annual and summer RE.

Statistics	Year-Round: January–...–December					Summer: June–July–August				
	Multipliers of the raw data									
	1976–1992	1993–2013	2014–2023	1976–1992	1993–2013	2014–2023				
	CC1	CC2	CC1	CC2	CC=1	CC1	CC2	CC1	CC2	CC=1
Mod1										
<del>MREMRD</del>	–3.3	–0.9	–4.2	–2.6	–0.2	–1.9	–2.4	–3.5	–1.9	–0.5
<del>MAEMA</del>	3.9	2.5	4.7	3.2	1.0	4.0	3.4	4.4	2.8	1.8
<del>RMSEMRD</del>	4.4	2.9	5.0	3.5	1.2	4.4	4.5	4.8	3.4	2.6
R	0.82	0.86	0.77	0.83	0.93	0.92	0.93	0.57	0.65	0.96
SD	3.2	3.0	2.8	2.4	1.4	4.2	4.0	3.3	2.9	2.7
Mod2										
<del>MREMRD</del>	–0.9	–1.0	–0.5	–0.6	–0.3	–1.0	–1.1	–0.6	–0.6	–0.0
<del>MAEMA</del>	2.1	2.0	2.0	1.8	0.6	3.3	2.9	2.2	1.9	1.3
<del>RMSEMRD</del>	2.6	2.4	2.7	2.3	0.8	4.0	3.7	2.8	2.4	1.8
R	0.90	0.92	0.81	0.86	0.97	0.93	0.94	0.72	0.79	0.98
SD	2.6	2.4	2.7	2.3	0.8	4.1	3.7	2.8	2.4	1.9
Mod3										
<del>MREMRD</del>	–0.4	–0.5	–0.9	–0.9	–0.6	–0.2	–0.1	–0.3	–0.3	–0.1

<u>MAEMA</u>	1.5	1.4	2.7	2.8	0.8	3.0	3.2	2.9	2.9	2.1
<u>D</u>										
<u>RMSE</u>	1.7	1.9	3.4	3.6	0.9	3.7	3.7	3.7	3.7	2.5
<u>SD</u>										
R	0.96	0.94	0.70	0.67	0.97	0.94	0.94	0.50	0.43	0.92
SD	1.8	2.0	3.4	3.5	0.8	3.9	3.9	3.8	3.8	2.7

### 3.3 Trend analyses

#### 3.3.1 The erythema annual and summer radiant exposures in the period 1976-2023

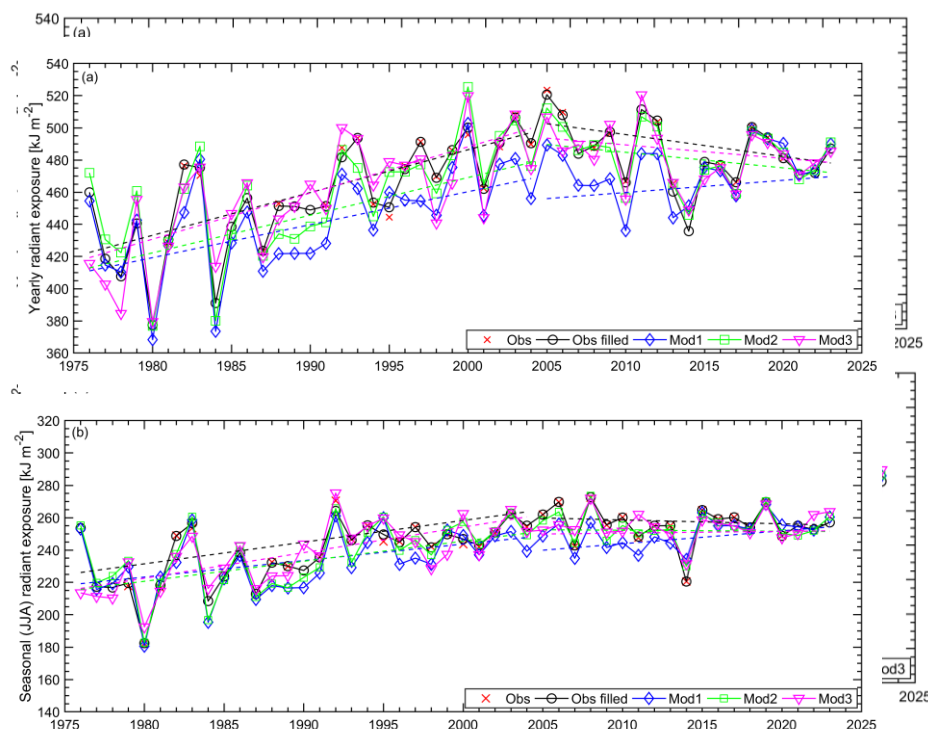
Trend analyses are applied to the erythema annual and summer RE based on daily ~~(for re-evaluated RE. There are~~  
~~two series to be considered when dealing with the measured data. The first, labelled OBS, uses only re-evaluated~~  
~~observations with filled gaps, OBS<sub>F</sub>, and Mod1) and the monthly RE (for Mod2 and Mod3). Gaps in the~~  
~~measurements were filled using Mod1 average is calculated when at least 14 daily ERE values are available. The~~  
~~second one, OBS<sub>F</sub>, contains all the daily gaps filled by Mod 1 simulations. Two versions of the OBS<sub>F</sub>, Mod2 and~~  
~~Mod3 time series are possible because of the use of CC1 and CC2 multipliers on For other models used there are~~  
~~no gaps. In case of Mod1, erythema annual and summer RE are built using the raw (1976–2013) re-evaluated daily~~  
~~measurements ERE values. For Mod2 and Mod3, the Mod1 time series, only one series was available for analysis,~~  
~~as this model monthly reconstructed erythema RE using values are summed over the proxy values and the model~~  
~~coefficients estimated from the KZ616 measurements (2014–2023), which did not require calibration, year and~~  
~~summer season.~~

~~For Mod2 and Mod3, two variants of the time series were available as these models required the 1976–1992 and~~  
~~1993–2013 mean values taken from the re-evaluated measurements with two possible options (CC1 or CC2) for~~  
~~the calibration multipliers. The 1976–2023 time series for the erythema annual and summer RE using CC1 and~~  
~~CC2 calibration correction multipliers are shown in Fig. 8 and Fig. A1, respectively. Fig.8a (Fig.A1a) and Fig.8b~~  
~~(Fig.A1b) are for the erythema annual (and summer) RE.~~

Linear regression lines are superimposed on the graphs to illustrate the long-term variability in the time series.  
Two independent lines are drawn to account for a change in the trend pattern observed in the time series somewhere  
in the early 2000s. The year of the trend change was calculated by examining the performance of fifteen  
combinations of this two-line pattern, varying the year of the trend change point (from 1995 to 2009). The best fit  
with maximum coefficients of determination ~~coefficients~~ was found for the trend change point in 2005. Therefore,  
the slopes of the regression lines (in kJ m<sup>-2</sup> per year) and the trend values (in % per year) shown in Table 56 and  
Table 67, respectively, are calculated for the 1976–2004 and 2005–2023 periods. Standard ~~deviations~~ errors of the  
trend estimates are calculated according to Eq. (15) ~~if the consecutive values accounting for the correction for the~~  
~~autocorrelation~~ in the trend residuals ~~are positively correlated, i.e. if~~ the autocorrelation coefficient with 1-yr lag,  
 $R_{k+1} > 0$ , is positive (also shown in Tables ~~4–5–6~~).



The interannual variations and trend lines of erythemal annual RE are close to each other when comparing the upper graphs in Fig. 8 and Fig.A1. This can also be observed for the summers when comparing the corresponding lower plots. At the beginning of the RB observations (1976–1986), there were large oscillations from year to year, suggesting an instrumental problem with the data. However, all modelled time series show quite similar oscillations for this period, supporting the ~~thesis~~hypothesis that a specific combination of TCO<sub>3</sub> and cloud transparency may be responsible for such oscillations.



**Figure 8.** Time series (1976–2023) of the erythemal radiant exposures from ~~reevaluated~~re-evaluated observations (Obs), ~~reevaluated~~re-evaluated observations with filled gaps (Obs filled), and model estimates (Mod1, Mod2, and Mod3) using the CC1 version of the ~~calibration~~correction coefficients: (a) annual (January–December) radiant exposures; (b) summer (June–Aug) radiant exposures. Dashed lines represent the linear trends calculated for the period 1976–2004 and 2005–2023.

The slopes of the linear fit to the analysed time series (Table 56) show a statistically significant positive trend between 1976 and 2004 of around 20–30 kJ m<sup>-2</sup> and 10–20 kJ m<sup>-2</sup> per decade in the annual and summer data, respectively. The trends are mostly insignificant for the period 2005–2023, with only one exception (for the Mod1 data) with a continued positive trend of ~ 10 kJ m<sup>-2</sup> per decade. The corresponding trend values expressed in dimensionless units (Table 67) have the same values of about 4–7% per 10 years in the former period for both the annual and summer time series. In the latter period, the positive trend of Mod1 is ~3 % per 10 years. ~~The smallest~~Mod 1 and Mod 3 (with CC2) gave the largest~~lowest and highest trends are always provided by Mod1 and Mod3 with CC2 calibration coefficients, respectively.~~ However, the differences between these trends are within the range of ± 2 standard errors of the trend estimates, taking into account the autocorrelation in the residuals of the models (column SE<sub>LS, COR</sub> in Table 56).



By averaging all available statistically significant annual and summer trend values shown in the third and seventh columns of Table 56 and Table 67, the following trends and their standard errors are obtained: for the period 1976–2004:  $27.4 \pm 4.4 \text{ kJ m}^{-2}$  and  $5.64 \pm 0.92 \%$  per decade for the erythral annual RE, and  $14.3 \pm 4.3 \text{ kJ m}^{-2}$  and  $5.63 \pm 1.03 \%$  per decade for the erythral summer RE. These values correspond to the average trend from the two series based only on the ~~reevaluated~~re-evaluated measurements ( $\text{OBS}_F$  values in the Tables), i.e.  $28.7 \text{ kJ m}^{-2}$  and  $5.9 \%$  per decade for the erythral annual RE, and  $14.3 \text{ kJ m}^{-2}$  and  $5.6 \%$  per decade for the erythral summer RE.

**Table 56.** Trends ( $\text{kJ m}^{-2}$  per year) by the linear least-squares fit to the time series of erythral annual and summer radiant exposures shown in Fig.8 and Fig.A1 calculated for the periods 1976–2004 and 2005–2023.  $SE_{\text{ELS, COR}}$  denotes the standard error of the trend estimate taking into account the autocorrelation (with a lag of 1 year) in the series of the residuals of the trend model.  $R_{k+1}$  denotes the correlation coefficient in the lagged residuals. Bold font indicates a statistically significant trend value at the 2-sigma level.

Data Type	Correct. Method	Annual (January...–December) sum [ $\text{kJ m}^{-2}$ ]				Summer (June–July–August) sum [ $\text{kJ m}^{-2}$ ]			
		Trends <sub>1976–2004</sub>		Trends <sub>2005–2023</sub>		Trends <sub>1976–2004</sub>		Trends <sub>2005–2023</sub>	
		Trend $\pm$ $SE_{\text{ELS, COR}}$	$R_{k+1}$	Trend $\pm$ $SE_{\text{ELS, COR}}$	$R_{k+1}$	Trend $\pm$ $SE_{\text{ELS, COR}}$	$R_{k+1}$	Trend $\pm$ $SE_{\text{ELS, COR}}$	$R_{k+1}$
$\text{OBS}_F$	CC1	<b><math>2.66 \pm 0.52</math></b>	–0.11	$-1.36 \pm 0.98$	0.17	<b><math>1.34 \pm 0.37</math></b>	0.08	$-0.24 \pm 0.49$	–0.32
	CC2	<b><math>3.08 \pm 0.52</math></b>	–0.06	$-0.45 \pm 0.87$	0.14	<b><math>1.52 \pm 0.37</math></b>	0.07	$-0.26 \pm 0.48$	–0.30
Mod1	–	<b><math>2.05 \pm 0.57</math></b>	–0.19	$0.76 \pm 0.78$	0.08	<b><math>1.02 \pm 0.38</math></b>	–0.13	<b><math>0.80 \pm 0.36</math></b>	–0.20
Mod2	CC1	<b><math>2.34 \pm 0.61</math></b>	–0.16	$-0.97 \pm 0.85$	0.12	<b><math>1.24 \pm 0.39</math></b>	–0.08	$-0.06 \pm 0.41$	–0.38
	CC2	<b><math>2.84 \pm 0.61</math></b>	–0.10	$-0.30 \pm 0.79$	0.10	<b><math>1.50 \pm 0.39</math></b>	–0.01	$0.29 \pm 0.41$	–0.32
Mod3	CC1	<b><math>2.84 \pm 0.56</math></b>	–0.21	$-0.84 \pm 0.76$	–0.08	<b><math>1.58 \pm 0.32</math></b>	0.02	$0.11 \pm 0.37$	–0.13
	CC2	<b><math>3.34 \pm 0.54</math></b>	–0.22	$-0.17 \pm 0.72$	–0.13	<b><math>1.82 \pm 0.20</math></b>	0.05	$0.46 \pm 0.36$	–0.13

**Table 67.** Same as Table 56, but the results are for the trend values expressed in % per year.

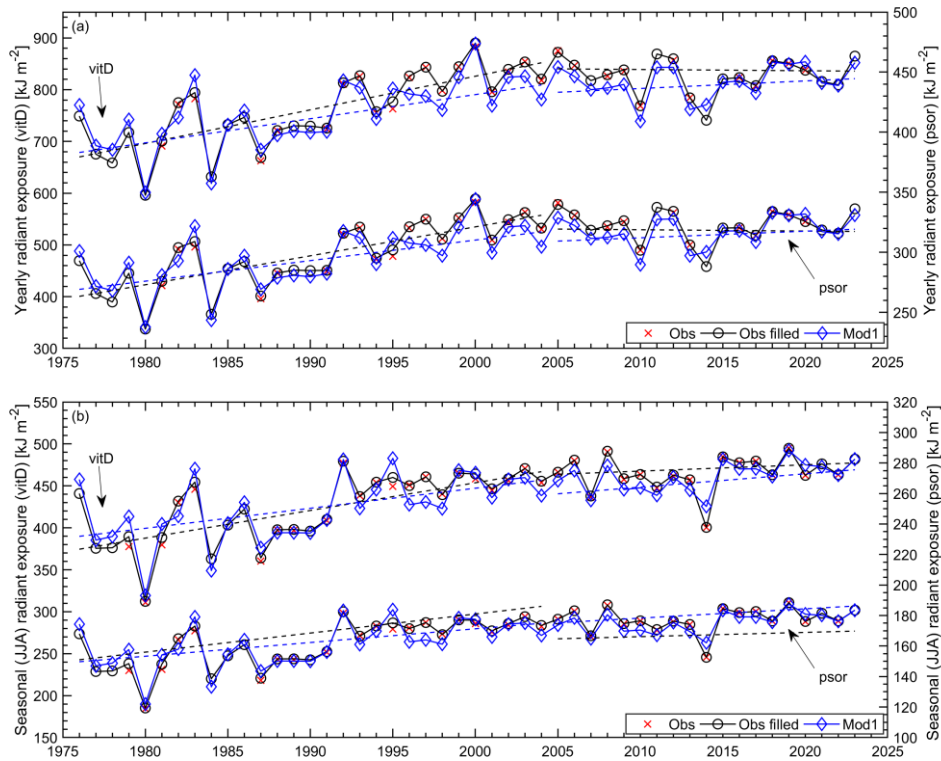
Data Type	Correct. Method	Annual (January...–December) sum [% yr <sup>–1</sup> ]				Summer (June–July–August) sum [% yr <sup>–1</sup> ]			
		Trends <sub>1976–2004</sub>		Trends <sub>2005–2023</sub>		Trends <sub>1976–2004</sub>		Trends <sub>2005–2023</sub>	
		Trend $\pm$ $SE_{\text{ELS, COR}}$	$R_{k+1}$	Trend $\pm$ $SE_{\text{ELS, COR}}$	$R_{k+1}$	Trend $\pm$ $SE_{\text{ELS, COR}}$	$R_{k+1}$	Trend $\pm$ $SE_{\text{ELS, COR}}$	$R_{k+1}$
$\text{OBS}_F$	CC1	<b><math>0.54 \pm 0.11</math></b>	–0.11	$-0.28 \pm 0.17$	0.17	<b><math>0.52 \pm 0.14</math></b>	0.08	$-0.09 \pm 0.19$	–0.32
	CC2	<b><math>0.64 \pm 0.11</math></b>	–0.06	$-0.09 \pm 0.16$	0.14	<b><math>0.60 \pm 0.14</math></b>	0.07	$0.10 \pm 0.19$	–0.30
Mod1	–	<b><math>0.42 \pm 0.12</math></b>	–0.19	$0.16 \pm 0.15$	0.08	<b><math>0.40 \pm 0.15</math></b>	–0.13	<b><math>0.31 \pm 0.14</math></b>	–0.20
Mod2	CC1	<b><math>0.48 \pm 0.13</math></b>	–0.16	$-0.20 \pm 0.15$	0.12	<b><math>0.49 \pm 0.15</math></b>	–0.08	$-0.02 \pm 0.16$	–0.38
	CC2	<b><math>0.59 \pm 0.12</math></b>	–0.10	$-0.06 \pm 0.15$	0.10	<b><math>0.59 \pm 0.15</math></b>	–0.01	$0.11 \pm 0.16$	–0.32
Mod3	CC1	<b><math>0.59 \pm 0.12</math></b>	–0.21	$-0.17 \pm 0.16$	–0.08	<b><math>0.62 \pm 0.12</math></b>	0.02	$0.04 \pm 0.15$	–0.13
	CC2	<b><math>0.69 \pm 0.11</math></b>	–0.22	$-0.04 \pm 0.15$	–0.13	<b><math>0.72 \pm 0.13</math></b>	0.05	$0.18 \pm 0.14$	–0.13

### 3.3.2 The vitamin D<sub>3</sub> and antipsoriatic annual and summer radiant exposures in the period 1976–2023

The ~~standard biometer~~commercial EBRs used to monitor erythral irradiance can also measure non-erythral irradiance (Czerwińska and Krzyścin, 2024a). Figure 5 and Table A1 provide that the daily vitamin D<sub>3</sub> and antipsoriatic RE derived from the KZ616 measurements agree with the directly measured BS64 values in the same way as the original (erythral) KZ616 data. This supports the method of the transfer from erythral irradiance to non-erythral irradiance proposed by Czerwińska and Krzyścin (2024a).

Figure 9 shows the time series of ~~the~~ annual and summer values of ~~the~~ previtamin D<sub>3</sub> synthesis and psoriasis healing RE from 1976 to 2023. It ~~looks like~~appears that these time series are very similar ~~when comparing the~~

vitamin D<sub>3</sub> to the antipsoriatic time series. Moreover, In addition, these time series are similar to the erythema series shown in Fig.8. The correlation coefficients between the pairs of time series shown in Fig.8 and Fig.9, i.e. erythema & vitamin D<sub>3</sub>, erythema & psoriasis, vitamin D<sub>3</sub> & psoriasis, ~~are were~~ in the range [0.90, >0.999] with the ~~smallest lowest~~ value for the ~~eases of~~ erythema & vitamin D<sub>3</sub>, ~~and~~ erythema & psoriasis ~~calculations using pairs~~ ~~when~~ the summer data from Mod1 simulations ~~were considered~~.



**Figure 9.** Time series (1976–2023) of the previtamin D<sub>3</sub> synthesis and psoriasis healing radiant exposures from ~~reevaluatedre-evaluated~~ observations (Obs), ~~reevaluatedre-evaluated~~ observations with filled gaps (Obs filled), and model Mod1 estimates (Mod1) using the CC2 version of the ~~calibrationcorrection~~ coefficients: (a) annual (January–December) ~~radiant~~ exposures; (b) summer (June–July–August) ~~radiant~~ exposures. Dashed lines represent the linear trends calculated for the period 1976–2004 and 2005–2023.

Table 78 shows the trend values for the period 1976–2004 and 2005–2023 from the time series calculated using the erythema DRE multiplied by the transfer coefficients defined by Eq. (4). The transfer coefficients depend on only two parameters (TCO<sub>3</sub> and SZA), even on cloudy days, as previously shown by Czerwińska and Krzyścin (2024a). The statistically significant trend values for previtamin D<sub>3</sub> synthesis and psoriasis clearance are slightly higher, by about 1–1.5 percentage points per decade, than the corresponding trend values for the erythema ~~appearance~~ shown in Table 67. Taking into account the standard error of the trend estimate of about 1% per decade, it cannot be said that the differences between the trends are statistically significant.

**Table 78.** Same as Table 67, but trend values are for previtamin D<sub>3</sub> synthesis and psoriasis clearance.

Data Type	Correct. Method	Annual (January–December) RE [% per year]				Summer (June–July–August) RE [% per year]			
		Trends <sub>1976–2004</sub>		Trends <sub>2005–2023</sub>		Trends <sub>1976–2004</sub>		Trends <sub>2005–2023</sub>	
		Trend ± SE <sub>LS</sub> , COR	R <sub>k+1</sub>	Trend ± SE <sub>LS</sub> , COR	R <sub>k+1</sub>	Trend ± SE <sub>LS</sub> , COR	R <sub>k+1</sub>	Trend ± SE <sub>LS</sub> , COR	R <sub>k+1</sub>

Previtamin D <sub>3</sub> synthesis									
OBS <sub>F</sub>	CC1	<b>0.70 ±0.12</b>	-0.12	-0.27 ±0.22	0.12	<b>0.64 ±0.16</b>	0.06	-0.07 ±0.20	-0.25
	CC2	<b>0.77 ±0.12</b>	-0.07	-0.03 ±0.19	0.08	<b>0.71 ±0.15</b>	0.05	0.16 ±0.19	-0.32
Mod1	–	<b>0.56 ±0.14</b>	-0.20	0.17 ±0.16	0.02	<b>0.51 ±0.16</b>	-0.15	<b>0.34 ±0.14</b>	-0.18
Psoriasis clearance									
OBS <sub>F</sub>	CC1	<b>0.66 ±0.12</b>	-0.13	-0.27 ±0.21	0.15	<b>0.63 ±0.15</b>	0.06	-0.07 ±0.20	-0.25
	CC2	<b>0.74 ±0.12</b>	-0.08	-0.03 ±0.20	0.10	<b>0.70 ±0.15</b>	0.05	0.16 ±0.19	-0.26
Mod1	–	<b>0.53 ±0.13</b>	-0.20	0.17 ±0.18	0.14	<b>0.51 ±0.16</b>	-0.15	<b>0.34 ±0.14</b>	-0.18

## 4 Summary and Discussion

One of the world's longest measurements of solar UV radiation at the Earth's surface (and probably the longest taken by erythral biometers) comes from Belsk. Measurements began in 1975 and continuous monitoring started on 1 January 1976. To the authors' knowledge, the longest UV monitoring series began in Moscow in 1968 with a broadband (300–380 nm) instrument developed at the Moscow State University Meteorological Observatory (Chubarova et al., 2000).

Several biometers participated in UV monitoring at Belsk, starting with RB, which operated until 31 December 1992. Subsequently, biometers SL501-A (#919 and #2011) and, since 5 August 2013, KZ616 have participated in UV monitoring. Each of these instruments has individual characteristics (spectral response, cosine error, ageing rate) and technical solutions, e.g. RB was not temperature stabilised and its output was in solar burn units. Therefore, a retrospective re-evaluation of the Belsk UV time series was necessary, and the homogenisation of the data from 1976 to 2023 is presented in this article.

Belsk is a unique observatory where UVUVR monitoring has been accompanied by monitoring of ozone (TCO<sub>3</sub>), aerosol optical properties (AOD) and cloud characteristics (sunshine duration, CIDCI from global solar irradiance measurements), i.e. basic input parameters to a radiative transfer model allowing reconstruction of the erythral RE. In addition, collocated BS64 measurements of UVUVR spectra allow monthly verification are used in the frequent (every month) checking of actual KZ616 performance. BS64 spectral measurements also allow assessment of the quality of Czerwińska and Krzyścin (2024a) retrieval to convert standard erythral measurements to the non-erythral BE irradiance (see the cases of the vitamin D<sub>3</sub> and antipsoriatic DRE in Figure 5).

Model simulations of erythral DRE and UVI under cloudless sky provide a basis for the correction procedure of raw UV data. A selection of clear sky conditions throughout the entire day from the daily proxy values (relative sunshine duration and RE from global solar irradiation), which were available for Belsk, is not straightforward as only the examination of the daily course of these measurements would allow to capture cloudless moments within the day. Therefore, two very different calibration configurations Two sets of raw UVR data multipliers (CC1 and CC2 as defined in section 2.3.2) have been proposed to assess the uncertainty range of the calibration correction method, applied to the raw UVR data. The reevaluated re-evaluated time series appear quite similar, i.e. the difference between these series is within a few percentage points (Fig.6, Table 4 and Table 5). There was no need for reevaluation of to re-evaluate the KS616 data for the period 2014–2023 as shown by the comparisons because they agreed well with the BS64 data (Fig.3 and Fig.-5).

Statistical Regression models trained on the KZ616 data for the period 2014–2023 allowed the data to be reconstructed from the beginning of UVUVR observations at Belsk. These reconstructed series allowed

independent examination of the pattern of interannual variability (which was unexpectedly large before 1985) and trends in the erythral annual and summer RE. The ~~statistical~~regression models generally mimic the observed long-term variability in the ~~reevaluated~~re-evaluated daily erythral exposures. The statistically significant trend of ~6 % per decade with a standard error of ~1 % per decade for the period 1976–2005 can be calculated (for both erythral annual and summer RE) by averaging trends from the sample of seven versions of trend estimates from ~~reevaluated~~re-evaluated and reconstructed data- (Table 7). All individual trend values are within the range of the mean trend  $\pm 2$  standard error (i.e. there is no outlier in this trend sample).

The standard errors for the individual trend estimates are in the range of 1-1.5% per decade, i.e. quite close to the standard error of the averaged trend derived from the trend sample. This supports the robustness of the trend estimates in annual and summer RE for the 1976–2005 parts of the Belsk time series. In addition, it also appears that the ~~very different calibration~~correction methods applied to the 1976–2013 raw ~~UV~~UVR data, based on the comparisons of clear-sky erythral DRE (CC1 method) and noon UVI (CC2 method), lead to differences in the individual 1976–2005 trend estimates of about 1 % per decade (see Table 67 for the trend differences between pairs of OBS<sub>F</sub>, Mod2 and Mod3 calculated with the CC1 and CC2 correction applied to the raw time series).

We found that ~~Mod1 could provide reasonable-our DRE estimates of DRE~~ for all biological effects considered (erythema, vitamin D<sub>3</sub> and psoriasis), ~~i.e.~~ were close to those obtained from the Brewer's spectra with a bias of ~~less than 2 ~ -1~~ % and a standard deviation of ~ 9 % (Table A1) for the part of the year when ~~UV radiation~~UVR is of particular interest, when the midday SZA is less than 45° (i.e. below the shadow length), according to the so-called shadow rule for protection against high ~~UV~~UVR (Downham, 1998).

Krzyścin et al. (2011) found a trend of 5.6 %  $\pm$  0.9 % (1 $\sigma$ ) per decade in the erythral annual RE for the period 1976–2008. This is in good agreement with the present trend estimate, regardless of the ~~very different calibration~~correction methods used. The correction of the SL501 A data carried out in 2011 was based on simultaneous measurements with KZ616 for the period 2008–2009 and further corrections for the instrument ageing using TUV cloudless sky simulations.

Similar trend estimates for erythral radiation can be inferred from the reconstructed erythral time series for the Moscow region based on the ~~UV~~UVR measurements by the broadband (300–380 nm) radiometer (Chubarova et al., 2018) and the statistically reconstructed erythral radiation series for Hradec Kralowe (Čížková et al., 2018). For the Moscow region, the authors reported a statistically significant positive trend of more than 5 % per decade for the period 1979-2015. Volpert and Chubarova (2021) revealed the decadal trend in the reconstructed erythral UV irradiance over the Moscow region for the warm season (May–September) of 5.1 %  $\pm$  1.1 % per decade in the period 1979–2016. Estimates from the smoothed pattern of annual erythral exposures taken from Fig. 2c by Čížková et al. (2018) for 1976 (~1.20 kJ m<sup>-2</sup> for the annual mean of erythral daily RE) and 2005 (~1.40 kJ m<sup>-2</sup>) give a trend of ~5% per decade for the period 1976–2004. From around 2005, both time series show a levelling off. Trends calculated here from the RE time series for other biological effects (previtamin D<sub>3</sub> synthesis and psoriasis lesion clearance), using an approach analogous to that used for the erythema data, show very similar trends.

## 5 Code and data availability

All data have been published as free access TXT files and are made available through PANGAEA repository at <https://doi.org/10.1594/PANGAEA.972139> (Krzyścin et al., 2024) and IG PAS Data Portal repository: [https://doi.org/10.25171/InstGeoph\\_PAS\\_IGData\\_Biologically\\_Effective\\_Solar\\_Radiation\\_Belsk\\_1976\\_2023](https://doi.org/10.25171/InstGeoph_PAS_IGData_Biologically_Effective_Solar_Radiation_Belsk_1976_2023) (Krzyścin, 2024). ERA5 data are publicly accessible at <https://cds.climate.copernicus.eu/datasets/reanalysis-era5-single-levels?tab=overview> (ERA5, 2024, 2025). MERRA-2 data are accessible at <https://doi.org/10.5067/Q9QMY5PBNV1T> (GMAO, 2024, 2025). Coefficients of the linear regression are calculated by Matlab function (Matlab R2018a) – *fitlm(x,y)*.

## 6 Conclusions

It is widely accepted that the use of overlapping measurement series from different instruments increases the reliability of results obtained from single time series analyses. Consequently, the inclusion of at least two different time series for analyses of the variability of a selected quantity over the entire measurement period is also beneficial for assessing data quality and establishing confidence in the results obtained. This is illustrated by the current data archived in the PANGAEA (Krzyścin et al., 2024) and IG PANPAS Data Portal (Krzyścin, 2024). The daily characteristics of BE radiation at Belsk allow the elaboration of scenarios of human outdoor activities to obtain maximum health benefits from sunbathing while minimising the risk of erythematous overexposure. The long-term variability of erythematous radiation calculated for Belsk corresponds to that previously recorded at distant stations in central/eastern Europe, making these scenarios applicable to wider areas.

## Appendix A

Table A1 presents descriptive statistics (defined in Sect. 2.4) of the relative differences between biologically effective DRE measured by the ~~BS64~~KZ616 and the ~~KZ616~~BS64,  $100\%(RE_{EFF,BS64KZ616} - RE_{EFF,KZ616BS64})/RE_{EFF,BS64}$ . The vitamin D<sub>3</sub> (VitD) and antipsoriatic (Psor) RE were reconstructed from the erythematous (Eryt) RE (Sect. 2.3.3), but the Brewer RE values were calculated using the daily integral of the measured spectral irradiance weighted by the action spectra (Fig.1).

**Table A1. Descriptive statistics of the 2014-2023 relative differences between the daily biologically effective radiant exposure with the ~~Brewer spectrophotometer #064~~ and the Kipp & Zonen erythema ~~biometer~~radiometer (UV-S-AE-T #30616) ~~at Belsk and the Brewer spectrophotometer #064~~ in percent of the Brewer data for the different midday SZA ranges (SZA<sub>N</sub>) ~~used in the Mod1 setup~~.**

Statistics	SZA <sub>N</sub> <45°			SZA <sub>N</sub> [45°, 60°]			SZA <sub>N</sub> ≥60°			All SZA <sub>N</sub>		
	Eryt	VitD	Psor	Eryt	VitD	Psor	Eryt	VitD	Psor	Eryt	VitD	Psor
<del>MREMR</del> <u>D</u>	-0.6	-1.5	-0.7	-2.5	-6.6	-3.3	-1.7	-13.2	-1.7	-1.4	-6.8	-1.6
<del>MAEMA</del> <u>D</u>	5.3	6.0	5.6	4.9	7.9	5.4	6.8	14.7	7.0	5.8	9.6	6.1
<del>RMSE%R</del> <u>MSD</u>	8.7	9.2	9.0	7.2	10.0	7.8	10.0	16.3	10.3	8.9	12.4	9.3
<del>SD%</del>	8.7	9.1	9.0	6.8	7.6	7.1	9.9	9.5	10.2	8.8	10.4	9.1

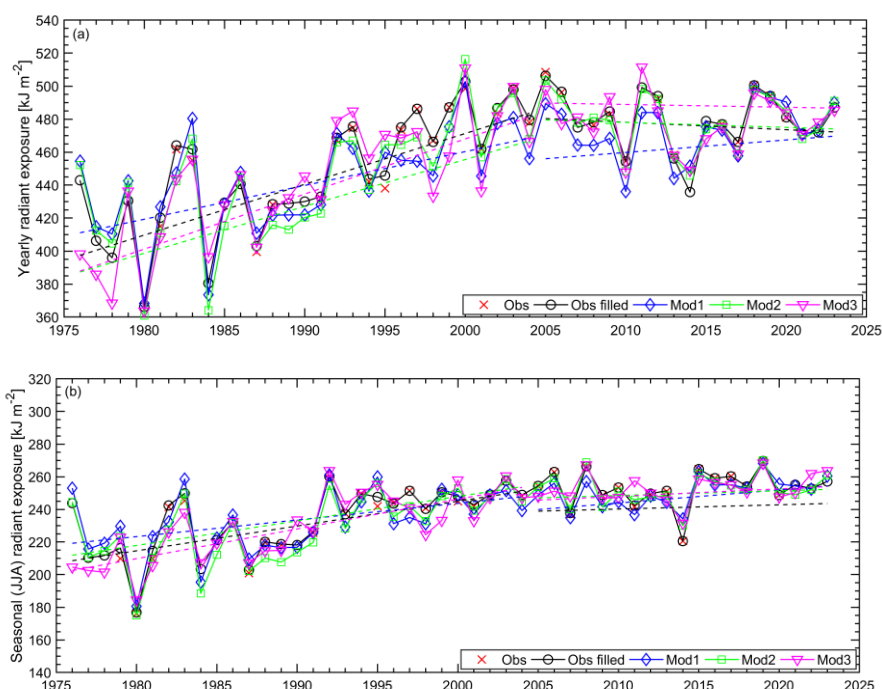


Figure A1. Same as Fig.8, but for the ~~reevaluated~~re-evaluated observations with the CC2 correction coefficients.

**Authors contributions.** Conceptualisation, JK and AC; methodology, JK, AC, JJ, PS, and BR; validation, AC and IP; visualisation, AC; writing (original draft preparation), JK and AC; writing (review and editing), JK, AC, PS, and IP; funding acquisition, JK and JJ. All authors have read and agreed to the published version of the paper.

**Competing interests.** The contact author has declared that none of the authors has any competing interests.

**Acknowledgments.** This work was partially financially by the Chief Inspectorate of Environment Protection, contract number GIOŚ/31/2023/DMŚ/NFOŚ.

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