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Global Carbon Budget 2024

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590 **1. Abstract**

591 Accurate assessment of anthropogenic carbon dioxide (CO<sub>2</sub>) emissions and their redistribution among the  
592 atmosphere, ocean, and terrestrial biosphere in a changing climate is critical to better understand the global  
593 carbon cycle, support the development of climate policies, and project future climate change. Here we describe  
594 and synthesise [datasets](#) and [methodologies](#) to quantify the five major components of the global carbon budget  
595 and their uncertainties. Fossil CO<sub>2</sub> emissions (E<sub>FOS</sub>) are based on energy statistics and cement production data,  
596 while emissions from land-use change (E<sub>LUC</sub>) are based on land-use and land-use change data and bookkeeping  
597 models. Atmospheric CO<sub>2</sub> concentration is measured directly, and its growth rate (G<sub>ATM</sub>) is computed from the  
598 annual changes in concentration. The ocean CO<sub>2</sub> sink (S<sub>OCEAN</sub>) is estimated with global ocean biogeochemistry  
599 models and observation-based *f*CO<sub>2</sub>-products. The terrestrial CO<sub>2</sub> sink (S<sub>LAND</sub>) is estimated with dynamic  
600 global vegetation models. Additional lines of evidence on land and ocean sinks are provided by atmospheric  
601 inversions, atmospheric oxygen measurements and Earth System Models. The [sum of all sources and sinks](#)  
602 [results in the](#) carbon budget imbalance (B<sub>IM</sub>), a measure of imperfect data and incomplete understanding of the  
603 contemporary carbon cycle. All uncertainties are reported as ±1σ.

604 For the year [2023](#), E<sub>FOS</sub> increased by [1.3%](#) relative to [2022](#), with fossil emissions at [10.1 ± 0.5 GtC yr<sup>-1</sup>](#) ([10.3 ±](#)  
605 [0.5 GtC yr<sup>-1</sup>](#) when the cement carbonation sink is not included), E<sub>LUC</sub> was [1.0 ± 0.7 GtC yr<sup>-1</sup>](#), for a total  
606 anthropogenic CO<sub>2</sub> emission (including the cement carbonation sink) of [11.1 ± 0.9 GtC yr<sup>-1</sup>](#) ([40.6 ± 3.2 GtCO<sub>2</sub>](#)  
607 [yr<sup>-1</sup>](#)). Also, for [2023](#), G<sub>ATM</sub> was [5.9 ± 0.2 GtC yr<sup>-1</sup>](#) ([2.79 ± 0.1 ppm yr<sup>-1</sup>](#)), S<sub>OCEAN</sub> was [2.9 ± 0.4 GtC yr<sup>-1</sup>](#) and  
608 S<sub>LAND</sub> was [2.3 ± 1.0 GtC yr<sup>-1</sup>](#), with a [near zero](#) B<sub>IM</sub> ([-0.02 GtC yr<sup>-1</sup>](#)). The global atmospheric CO<sub>2</sub> concentration  
609 averaged over [2023](#) reached [419.3 ± 0.1 ppm](#). Preliminary data for [2024](#), suggest an increase in E<sub>FOS</sub> relative to  
610 [2023](#) of [+0.8%](#) ([-0.3%](#) to [1.9%](#)) globally, and atmospheric CO<sub>2</sub> concentration [increased by 2.8 ppm](#) reaching  
611 [422.5 ppm](#), [52%](#) above pre-industrial level (around 278 ppm in 1750). Overall, the mean and trend in the  
612 components of the global carbon budget are consistently estimated over the period 1959–[2023](#), with a near-zero  
613 overall budget imbalance, although discrepancies of up to around 1 GtC yr<sup>-1</sup> persist for the representation of  
614 annual to semi-decadal variability in CO<sub>2</sub> fluxes. Comparison of estimates from multiple approaches and  
615 observations shows: (1) a persistent large uncertainty in the estimate of land-use changes emissions, (2) a low  
616 agreement between the different methods on the magnitude of the land CO<sub>2</sub> flux in the northern extra-tropics,  
617 and (3) a discrepancy between the different methods on the [mean](#) ocean sink.

618 This living data update documents changes in methods and [datasets](#) applied to this most-recent global carbon  
619 budget as well as evolving community understanding of the global carbon cycle. The data presented in this  
620 work are available at <https://doi.org/10.18160/GCP-2024> (Friedlingstein et al., [2024](#)).

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## 2. Executive Summary

Global fossil CO<sub>2</sub> emissions (including cement carbonation) are expected to further increase in 2024 by 0.8%. The 2023 emission increase was 0.4 GtC yr<sup>-1</sup> (0.5 GtCO<sub>2</sub> yr<sup>-1</sup>) relative to 2022, bringing 2023 fossil CO<sub>2</sub> emissions to 10.1 ± 0.5 GtC yr<sup>-1</sup> (36.8 ± 1.8 GtCO<sub>2</sub> yr<sup>-1</sup>). Preliminary estimates based on data available suggest fossil CO<sub>2</sub> emissions to increase further in 2024, by 0.8% relative to 2023 (-0.3% to 1.9%), bringing emissions to 10.2 GtC yr<sup>-1</sup> (37.4 GtCO<sub>2</sub> yr<sup>-1</sup>).<sup>1</sup>

Emissions from coal, oil and gas in 2024 are expected to be slightly above their 2023 levels (by 0.2%, 0.9% and 2.4% respectively). Regionally, fossil emissions in 2024 are expected to decrease by 3.8% in the European Union (reaching 0.7 GtC, 2.4 GtCO<sub>2</sub>), and by 0.6% in the United States (1.3 GtC, 4.9 GtCO<sub>2</sub>). Emissions in China are expected to increase in 2024 by 0.2%, reaching 3.3 GtC, (12.0 GtCO<sub>2</sub>). Fossil emissions are also expected to increase by 4.6% in India (0.9 GtC, 3.2 GtCO<sub>2</sub>) and by 1.1% for the rest of the world (4.0 GtC, 14.5 GtCO<sub>2</sub>) in 2024. Emissions from international aviation and shipping (IAS) are also expected to increase by 7.8% (0.3 GtC, 1.2 GtCO<sub>2</sub>) in 2024.

Fossil CO<sub>2</sub> emissions decreased significantly in 22 countries with significantly growing economies during the decade 2014-2023. Altogether, these 22 countries contribute about 2.2 GtC yr<sup>-1</sup> (8.1 GtCO<sub>2</sub>) fossil fuel CO<sub>2</sub> emissions over the last decade, representing about 23% of world CO<sub>2</sub> fossil emissions.

Global CO<sub>2</sub> emissions from land-use, land-use change, and forestry (LULUCF) averaged 1.1 ± 0.7 GtC yr<sup>-1</sup> (4.1 ± 2.6 GtCO<sub>2</sub> yr<sup>-1</sup>) for the 2014-2023 period with a similar preliminary projection for 2024 of 1.1 ± 0.7 GtC yr<sup>-1</sup> (4.2 ± 2.6 GtCO<sub>2</sub> yr<sup>-1</sup>). Since the late-1990s, emissions from LULUCF show a statistically significant decrease at a rate of around 0.2 GtC per decade. Emissions from deforestation, the main driver of global gross sources, remain high at around 1.7 GtC yr<sup>-1</sup> over the 2014-2023 period, highlighting the strong potential of halting deforestation for emissions reductions. Sequestration of 1.2 GtC yr<sup>-1</sup> through re-forestation and forestry offsets two third of the deforestation emissions. Further, smaller emissions are due to other land-use transitions and peat drainage and peat fire. The highest emitters during 2014-2023 in descending order were Brazil, Indonesia, and the Democratic Republic of the Congo, with these 3 countries contributing more than half of global land-use CO<sub>2</sub> emissions.

Total anthropogenic emissions (fossil and LULUCF, including the carbonation sink) were 11.1 GtC yr<sup>-1</sup> (40.6 GtCO<sub>2</sub> yr<sup>-1</sup>) in 2023, with a marginally higher preliminary estimate of 11.3 GtC yr<sup>-1</sup> (41.6 GtCO<sub>2</sub> yr<sup>-1</sup>) for 2024. Total anthropogenic emissions have been stable over the last decade (zero growth rate over the 2014-2023 period), much slower than over the previous decade (2004-2013) with an average growth rate of 2.0% yr<sup>-1</sup>.

The remaining carbon budget for a 50% likelihood to limit global warming to 1.5°C, 1.7°C and 2°C above the 1850-1900 level has respectively been reduced to 65 GtC (235 GtCO<sub>2</sub>), 160 GtC (585 GtCO<sub>2</sub>) and 305

Deleted: 2023, to 1.4% above their pre-COVID-19 pandemic 2019 level...024 by 0.8%. The 2022...023 emission increase was 0.09...4 GtC yr<sup>-1</sup> (0.33... GtCO<sub>2</sub> yr<sup>-1</sup>) relative to 2021...022, bringing 2022...023 fossil CO<sub>2</sub> emissions to 9.9...0.1 ± 0.5 GtC yr<sup>-1</sup> (36.4... ± 1.8 GtCO<sub>2</sub> yr<sup>-1</sup>), virtually equal to the emissions level of 2019... Preliminary estimates based on data available suggest fossil CO<sub>2</sub> emissions to increase further in 2023...024, by 1.1...8% relative to 2022 (0...023 (-0.3% to 2...9%), bringing emissions to 10.0... GtC yr<sup>-1</sup> (36.8...7.4 GtCO<sub>2</sub> yr<sup>-1</sup>), 1.4% above the 2019 level. (... [20])

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<sup>1</sup> All 2024 growth rates use a leap year adjustment that corrects for the extra day in 2024.

820 GtC (1110 GtCO<sub>2</sub>) from the beginning of 2025, equivalent to around 6, 14 and 27 years, assuming 2024  
821 emissions levels.

822 The concentration of CO<sub>2</sub> in the atmosphere is set to reach 422.5 ppm in 2024, 52% above pre-industrial  
823 levels. The atmospheric CO<sub>2</sub> growth was  $5.2 \pm 0.02$  GtC yr<sup>-1</sup> (2.5 ppm) during the decade 2014-2023 (48% of  
824 total CO<sub>2</sub> emissions) with a preliminary 2024 growth rate estimate of around 5.9 GtC (2.8 ppm).

825 The ocean CO<sub>2</sub> sink has been stagnant since 2016 after rapid growth during 2002-2016, largely in  
826 response to large inter-annual climate variability. The ocean CO<sub>2</sub> sink was  $2.9 \pm 0.4$  GtC yr<sup>-1</sup> during the  
827 decade 2014-2023 (26% of total CO<sub>2</sub> emissions). A slightly higher value of 3.0 GtC yr<sup>-1</sup> is preliminarily  
828 estimated for 2024, which marks an increase in the sink since 2023 due to the prevailing El Niño and neutral  
829 conditions in 2024.

830 The land CO<sub>2</sub> sink continued to increase during the 2014-2023 period primarily in response to increased  
831 atmospheric CO<sub>2</sub>, albeit with large interannual variability. The land CO<sub>2</sub> sink was  $3.2 \pm 0.9$  GtC yr<sup>-1</sup> during  
832 the 2014-2023 decade (30% of total CO<sub>2</sub> emissions). The land sink in 2023 was  $2.3 \pm 1$  GtC yr<sup>-1</sup>, 1.6 GtC lower  
833 than in 2022, and the lowest estimate since 2015. This reduced sink is primarily driven by a response of tropical  
834 land ecosystems to the onset of the 2023-2024 El Niño event, combined with large wildfires in Canada in 2023.  
835 The preliminary 2024 estimate is around 3.2 GtC yr<sup>-1</sup>, similar to the decadal average, consistent with a land sink  
836 emerging from the El Niño state.

837 So far in 2024, global fire CO<sub>2</sub> emissions have been 11-32% higher than the 2014-2023 average due to  
838 high fire activity in both North and South America, reaching 1.6-2.2 GtC during January-September. In  
839 Canada, emissions through September were 0.2-0.3 GtC yr<sup>-1</sup>, down from 0.5-0.8 GtC yr<sup>-1</sup> in 2023 but still more  
840 than twice the 2014-2023 average. In Brazil, fires through September emitted 0.2-0.3 GtC yr<sup>-1</sup>, 91-118% above  
841 the 2014-2023 average due to intense drought. These fire emissions estimates should not be directly compared  
842 with the land use emissions or the land sink, because they represent a gross carbon flux to the atmosphere and  
843 do not account for post-fire recovery or distinguish between natural, climate-driven, and land-use-related fires.

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pre-industrial levels. The atmospheric CO<sub>2</sub> growth was  $5.2 \pm$   
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after rapid growth in the past two decades after low or no  
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annual climate variability. The estimates based on fCO<sub>2</sub>-  
products and models diverge with the growth of the ocean  
CO<sub>2</sub> sink in the past decade being a factor 2.5 larger than in  
the models. This discrepancy in the trend originates from all  
latitudes but is largest in the Southern Ocean. ...he ocean  
sink was  $2.9 \pm 0.4$  GtC yr<sup>-1</sup> during the decade 2013-  
2022...014-2023 (26% of total CO<sub>2</sub> emissions), and did not  
grow since 2019 due to a triple La Niña event.... A  
similar...lightly higher value of  $2.9 \pm 0.4$  GtC yr<sup>-1</sup> is  
preliminarily estimated for 2023...024, which marks an  
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due to the transition from La Niña to...revailing El Niño and  
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decade (2000-2009), with a preliminary 2023 estimate of  
around 2.9 GtC yr<sup>-1</sup>, significantly .... The land sink in 2023  
was  $2.3 \pm 1$  GtC yr<sup>-1</sup>, 1.6 GtC lower than in 2022, and  
attributed to the ...he lowest estimate since 2015. This  
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America, reaching 1.6-2.2 GtC during January- (... [31])

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## 949 1 Introduction

950 The concentration of carbon dioxide (CO<sub>2</sub>) in the atmosphere has increased from approximately 278 parts per  
 951 million (ppm) in 1750 (Gulev et al., 2021), the beginning of the Industrial Era, to  $419.3 \pm 0.1$  ppm in 2023 (Lan  
 952 et al., 2024; Figure 1). The atmospheric CO<sub>2</sub> increase above pre-industrial levels was, initially, primarily caused  
 953 by the release of carbon to the atmosphere from deforestation and other land-use change activities (Canadell et  
 954 al., 2021). While emissions from fossil fuels started before the Industrial Era, they became the dominant source  
 955 of anthropogenic emissions to the atmosphere from around 1950 and their relative share has continued to  
 956 increase until present. Anthropogenic emissions occur on top of an active natural carbon cycle that circulates  
 957 carbon between the reservoirs of the atmosphere, ocean, and terrestrial biosphere on time scales from sub-daily  
 958 to millennial, while exchanges with geologic reservoirs occur on longer timescales (Archer et al., 2009).

959 The global carbon budget (GCB) presented here refers to the mean, variations, and trends in the perturbation of  
 960 CO<sub>2</sub> in the environment, referenced to the beginning of the Industrial Era (defined here as 1750). This paper  
 961 describes the components of the global carbon cycle over the historical period with a stronger focus on the  
 962 recent period (since 1958, onset of robust atmospheric CO<sub>2</sub> measurements), the last decade (2014-2023), the last  
 963 year (2023) and the current year (2024). Finally, it provides cumulative emissions from fossil fuels and land-use  
 964 change since the year 1750, and since the year 1850 (the reference year for historical simulations in IPCC AR6)  
 965 (Eyring et al., 2016).

966 We quantify the input of CO<sub>2</sub> to the atmosphere by emissions from human activities, the growth rate of  
 967 atmospheric CO<sub>2</sub> concentration, and the resulting changes in the storage of carbon in the land and ocean  
 968 reservoirs in response to increasing atmospheric CO<sub>2</sub> levels, climate change and variability, and other  
 969 anthropogenic and natural changes (Figure 2). An understanding of this perturbation budget over time and the  
 970 underlying variability and trends of the natural carbon cycle is necessary to understand the response of natural  
 971 sinks to changes in climate, CO<sub>2</sub> and land-use change drivers, and to quantify emissions compatible with a given  
 972 climate stabilisation target.

973 The components of the CO<sub>2</sub> budget that are reported annually in this paper include separate and independent  
 974 estimates for the CO<sub>2</sub> emissions from (1) fossil fuel combustion and oxidation from all energy and industrial  
 975 processes; also including cement production and carbonation ( $E_{\text{FOS}}$ ; GtC yr<sup>-1</sup>) and (2) the emissions resulting  
 976 from deliberate human activities on land, including those leading to land-use change ( $E_{\text{LUC}}$ ; GtC yr<sup>-1</sup>); and their  
 977 partitioning among (3) the growth rate of atmospheric CO<sub>2</sub> concentration ( $G_{\text{ATM}}$ ; GtC yr<sup>-1</sup>), and the uptake of  
 978 CO<sub>2</sub> (the ‘CO<sub>2</sub> sinks’) in (4) the ocean ( $S_{\text{OCEAN}}$ ; GtC yr<sup>-1</sup>) and (5) on land ( $S_{\text{LAND}}$ ; GtC yr<sup>-1</sup>). The CO<sub>2</sub> sinks as  
 979 defined here conceptually include the response of the land (including inland waters and estuaries) and ocean  
 980 (including coastal and marginal seas) to elevated CO<sub>2</sub> and changes in climate and other environmental  
 981 conditions, although in practice not all processes are fully accounted for (see Section 2.10). Global emissions  
 982 and their partitioning among the atmosphere, ocean and land are in balance in the real world. Due to the  
 983 combination of imperfect spatial and/or temporal data coverage, errors in each estimate, and smaller terms not

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included in our budget estimate (discussed in Section 2.10), the independent estimates (1) to (5) above do not necessarily add up to zero. We hence estimate a budget imbalance ( $B_{IM}$ ), which is a measure of the mismatch between the estimated emissions and the estimated changes in the atmosphere, land and ocean, as follows:

$$B_{IM} = E_{FOS} + E_{LUC} - (G_{ATM} + S_{OCEAN} + S_{LAND}) \quad (1)$$

$G_{ATM}$  is usually reported in ppm yr<sup>-1</sup>, which we convert to units of carbon mass per year, GtC yr<sup>-1</sup>, using 1 ppm = 2.124 GtC (Ballantyne et al., 2012; Table 1). Units of gigatonnes of CO<sub>2</sub> (or billion tonnes of CO<sub>2</sub>) used in policy are equal to 3.664 multiplied by the value in units of GtC.

We also assess a set of additional lines of evidence derived from global atmospheric inversion system results (Section 2.7), observed changes in oxygen concentration (Section 2.8) and Earth System Models (ESMs) simulations (Section 2.9), all of these methods closing the global carbon balance (zero  $B_{IM}$ ).

We further quantify  $E_{FOS}$  and  $E_{LUC}$  by country, including both territorial and consumption-based accounting for  $E_{FOS}$  (see Section 2), and discuss missing terms from sources other than the combustion of fossil fuels (see Section 2.10, Supplement S1 and S2). We also assess carbon dioxide removal (CDR) (see Sect. 2.2 and 2.3). Land-based CDR is significant, but already accounted for in  $E_{LUC}$  in equation (1) (Sect 3.2.2). Other CDR methods, not based on vegetation, are currently several orders of magnitude smaller than the other components of the budget (Sect. 3.3), hence these are not included in equation (1), or in the global carbon budget tables or figures (with the exception of Figure 2 where CDR is shown primarily for illustrative purpose).

The global CO<sub>2</sub> budget has been assessed by the Intergovernmental Panel on Climate Change (IPCC) in all assessment reports (Prentice et al., 2001; Schimel et al., 1995; Watson et al., 1990; Denman et al., 2007; Ciais et al., 2013; Canadell et al., 2021), and by others (e.g. Ballantyne et al., 2012). The Global Carbon Project (GCP, [www.globalcarbonproject.org](http://www.globalcarbonproject.org), last access: 28 October 2024) has coordinated this cooperative community effort for the annual publication of global carbon budgets for the year 2005 (Raupach et al., 2007; including fossil emissions only), year 2006 (Canadell et al., 2007), year 2007 (GCP, 2008), year 2008 (Le Quéré et al., 2009), year 2009 (Friedlingstein et al., 2010), year 2010 (Peters et al., 2012a), year 2012 (Le Quéré et al., 2013; Peters et al., 2013), year 2013 (Le Quéré et al., 2014), year 2014 (Le Quéré et al., 2015a; Friedlingstein et al., 2014), year 2015 (Jackson et al., 2016; Le Quéré et al., 2015b), year 2016 (Le Quéré et al., 2016), year 2017 (Le Quéré et al., 2018a; Peters et al., 2017a), year 2018 (Le Quéré et al., 2018b; Jackson et al., 2018), year 2019 (Friedlingstein et al., 2019; Jackson et al., 2019; Peters et al., 2020), year 2020 (Friedlingstein et al., 2020; Le Quéré et al., 2021), year 2021 (Friedlingstein et al., 2022a; Jackson et al., 2022), year 2022 (Friedlingstein et al., 2022b), and most recently the year 2023 (Friedlingstein et al., 2023). Each of these papers updated previous estimates with the latest available information for the entire time series.

We adopt a range of  $\pm 1$  standard deviation ( $\sigma$ ) to report the uncertainties in our global estimates, representing a likelihood of 68% that the true value will be within the provided range if the errors have a gaussian distribution, and no bias is assumed. This choice reflects the difficulty of characterising the uncertainty in the CO<sub>2</sub> fluxes between the atmosphere and the ocean and land reservoirs individually, particularly on an annual basis, as well

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**Moved up [6]:** Units of gigatonnes of CO<sub>2</sub> (or billion tonnes of CO<sub>2</sub>) used in policy are equal to 3.664 multiplied by the value in units of GtC.<sup>¶</sup>

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1046 as the difficulty of updating the CO<sub>2</sub> emissions from land-use change. A likelihood of 68% provides an  
1047 indication of our current capability to quantify each term and its uncertainty given the available information.  
1048 The uncertainties reported here combine statistical analysis of the underlying data, assessments of uncertainties  
1049 in the generation of the [datasets](#), and expert judgement of the likelihood of results lying outside this range. The  
1050 limitations of current information are discussed in the paper and have been examined in detail elsewhere  
1051 (Ballantyne et al., 2015; Zscheischler et al., 2017). We also use a qualitative assessment of confidence level to  
1052 characterise the annual estimates from each term based on the type, amount, quality, and consistency of the  
1053 different lines of evidence as defined by the IPCC (Stocker et al., 2013).

1054 This paper provides a detailed description of the [datasets](#) and methodology used to compute the global carbon  
1055 budget estimates for the industrial period, from 1750 to [2024](#), and in more detail for the period since 1959. This  
1056 paper is updated every year using the format of 'living data' to keep a record of budget versions and the changes  
1057 in new data, revision of data, and changes in methodology that lead to changes in estimates of the carbon  
1058 budget. Additional materials associated with the release of each new version will be posted at the Global Carbon  
1059 Project (GCP) website (<http://www.globalcarbonproject.org/carbonbudget>, last access: [28 October 2024](#)), with  
1060 fossil fuel emissions also available through the Global Carbon Atlas (<http://www.globalcarbonatlas.org>, last  
1061 access: [28 October 2024](#)). All underlying data used to produce the budget can also be found at  
1062 <https://globalcarbonbudget.org/> (last access: [28 October 2024](#)). With this approach, we aim to provide the  
1063 highest transparency and traceability in the reporting of CO<sub>2</sub>, the key driver of climate change.

## 1064 2 Methods

1065 Multiple organisations and research groups around the world generated the original measurements and data used  
1066 to complete the global carbon budget. The effort presented here is thus mainly one of synthesis, where results  
1067 from individual groups are collated, analysed, and evaluated for consistency. We facilitate access to original  
1068 data with the understanding that primary [datasets](#) will be referenced in future work (see Table 2 for how to cite  
1069 the [datasets](#), and Section on data availability). Descriptions of the measurements, models, and methodologies  
1070 follow below, with more detailed descriptions of each component provided as Supplementary Information (S1 to  
1071 S5).

1072 This is the [19<sup>th</sup>](#) version of the global carbon budget and the [13<sup>th</sup>](#) revised version in the format of a living data  
1073 update in Earth System Science Data. It builds on the latest published global carbon budget of Friedlingstein et  
1074 al. ([2023](#)). The main changes this year are: the inclusion of (1) data to year [2023](#) and a projection for the global  
1075 carbon budget for year [2024](#); and (2) [an estimate of the 2024 projection of fossil emissions from Carbon](#)  
1076 [Monitor](#). Other methodological differences between recent annual carbon budgets ([2020](#) to [2024](#)) are  
1077 summarised in Table 3 and previous changes since 2006 are provided in Table [S9](#).

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1100 2.1 Fossil CO<sub>2</sub> emissions (E<sub>FOS</sub>)

1101 2.1.1 Historical period 1850-2023

1102 The estimates of global and national fossil CO<sub>2</sub> emissions (E<sub>FOS</sub>) include the oxidation of fossil fuels through  
1103 both combustion (e.g., transport, heating) and chemical oxidation (e.g. carbon anode decomposition in  
1104 aluminium refining) activities, and the decomposition of carbonates in industrial processes (e.g. the production  
1105 of cement). We also include CO<sub>2</sub> uptake from the cement carbonation process. Several emissions sources are not  
1106 estimated or not fully covered: coverage of emissions from lime production are not global, and decomposition of  
1107 carbonates in glass and ceramic production are included only for the “Annex 1” countries of the United Nations  
1108 Framework Convention on Climate Change (UNFCCC) for lack of activity data. These omissions are  
1109 considered to be minor. Short-cycle carbon emissions - for example from combustion of biomass - are not  
1110 included here but are accounted for in the CO<sub>2</sub> emissions from land use (see Section 2.2).

1111 Our estimates of fossil CO<sub>2</sub> emissions rely on data collection by many other parties. Our goal is to produce the  
1112 best estimate of this flux, and we therefore use a prioritisation framework to combine data from different  
1113 sources that have used different methods, while being careful to avoid double counting and undercounting of  
1114 emissions sources. The CDIAC-FF emissions dataset, derived largely from UN energy data, forms the  
1115 foundation, and we extend emissions to 2023 using energy growth rates reported by the Energy Institute (a  
1116 dataset formerly produced by BP). We then proceed to replace estimates using data from what we consider to be  
1117 superior sources, for example Annex 1 countries’ official submissions to the UNFCCC. All data points are  
1118 potentially subject to revision, not just the latest year. For full details see Andrew and Peters (2024).

1119 Other estimates of global fossil CO<sub>2</sub> emissions exist, and these are compared by Andrew (2020a). The most  
1120 common reason for differences in estimates of global fossil CO<sub>2</sub> emissions is a difference in which emissions  
1121 sources are included in the datasets. Datasets such as those published by the Energy Institute, the US Energy  
1122 Information Administration, and the International Energy Agency’s ‘CO<sub>2</sub> emissions from fuel combustion’ are  
1123 all generally limited to emissions from combustion of fossil fuels. In contrast, datasets such as PRIMAP-hist,  
1124 CEDS, EDGAR, and GCP’s dataset aim to include all sources of fossil CO<sub>2</sub> emissions. See Andrew (2020a) for  
1125 detailed comparisons and discussion.

1126 Cement absorbs CO<sub>2</sub> from the atmosphere over its lifetime, a process known as ‘cement carbonation’. We  
1127 estimate this CO<sub>2</sub> sink, from 1931 onwards, as the average of two studies in the literature (Cao et al., 2020; Guo  
1128 et al., 2021). Both studies use the same model, developed by Xi et al. (2016), with different parameterisations  
1129 and input data, with the estimate of Guo and colleagues being a revision of Xi et al. (2016). The trends of the  
1130 two studies are very similar. Since carbonation is a function of both current and previous cement production, we  
1131 extend these estimates to 2023 by using the growth rate derived from the smoothed cement emissions (10-year  
1132 smoothing) fitted to the carbonation data. In the present budget, we always include the cement carbonation  
1133 carbon sink in the fossil CO<sub>2</sub> emission component (E<sub>FOS</sub>).

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1141 We use the Kaya Identity for a simple decomposition of CO<sub>2</sub> emissions into the key drivers (Raupach et al.,  
1142 2007). While there are variations (Peters et al., 2017a), we focus here on a decomposition of CO<sub>2</sub> emissions into  
1143 population, GDP per person, energy use per GDP, and CO<sub>2</sub> emissions per energy. Multiplying these individual  
1144 components together returns the CO<sub>2</sub> emissions. Using the decomposition, it is possible to attribute the change  
1145 in CO<sub>2</sub> emissions to the change in each of the drivers. This method gives a first-order understanding of what  
1146 causes CO<sub>2</sub> emissions to change each year.

147 **2.1.2 2024 projection**

148 We provide a projection of global fossil CO<sub>2</sub> emissions in 2024 by combining separate projections for China,  
149 USA, EU, India, and for all other countries combined. The methods are different for each of these. For China we  
150 combine monthly fossil fuel production data from the National Bureau of Statistics and trade data from the  
151 Customs Administration, giving us partial data for the growth rates to date of natural gas, petroleum, and  
152 cement, and of the apparent consumption itself for raw coal. We then use a regression model to project full-year  
153 emissions based on historical observations. For the USA our projection is taken directly from the Energy  
154 Information Administration's (EIA) Short-Term Energy Outlook (EIA, 2024), combined with the year-to-date  
155 growth rate of cement clinker production. For the EU we use monthly energy data from Eurostat to derive  
156 estimates of monthly CO<sub>2</sub> emissions through July, with coal emissions extended through September using a  
157 statistical relationship with reported electricity generation from coal and other factors. For natural gas we use  
158 Holt-Winters to project the last four months of the year. EU emissions from oil are derived using the EIA's  
159 projection of oil consumption for Europe. EU cement emissions are based on available year-to-date data from  
160 three of the largest producers, Germany, Poland, and Spain. India's projected emissions are derived from  
161 estimates through August (July for coal) using the methods of Andrew (2020b) and extrapolated assuming  
162 seasonal patterns from before 2019. Emissions from international transportation (bunkers) are estimated  
163 separately for aviation and shipping. Changes in aviation emissions are derived primarily from OECD monthly  
164 estimates, extrapolated using the growth rates of global flight miles from Airportia, and then the final months  
165 are projected assuming normal patterns from previous years. Changes in shipping emissions are derived from  
166 OECD monthly estimates for global shipping. Emissions for the rest of the world are derived for coal and  
167 cement using projected growth in economic production from the IMF (2023) combined with extrapolated  
168 changes in emissions intensity of economic production, for oil using a global constraint from EIA; and for  
169 natural gas using a global constraint from IEA. More details on the E<sub>FOS</sub> methodology and its 2024 projection  
170 can be found in Supplement S.1.

171 For the first time this year, we cross check our 2024 projection with a 2024 projection from Carbon Monitor.  
172 Carbon Monitor is an open access dataset (<https://carbonmonitor.org/>) of daily emissions constructed using  
173 hourly to daily proxy data (e.g., electricity consumption, travel patterns, etc) instead of energy use data.  
174 Available Carbon Monitor estimated emissions from January to August are combined to a new projection for  
175 September to December to give a full year 2024 estimate. The September to December projections are estimated  
176 by leveraging seasonal patterns from 2019-2023 daily CO<sub>2</sub> emission data from Carbon Monitor. A regression  
177 model is applied separately for individual countries to obtain their respective 4-month forecast. First, the

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seasonality component for each month is assessed based on daily average emissions from 2019 to 2023, excluding 2020 due to the COVID-19 pandemic. Then, a linear regression model is constructed using the calculated seasonal components and the daily average emissions for the months from January to August 2024. The resulting model is used to project carbon emissions for the remaining months of 2024. The uncertainty range is calculated by using historical monthly variance of seasonal components.

2.2 CO<sub>2</sub> emissions from land-use, land-use change and forestry (E<sub>LUC</sub>)

2.2.1 Historical period 1850-2023

The net CO<sub>2</sub> flux from land-use, land-use change and forestry (E<sub>LUC</sub>, called land-use change emissions in the rest of the text) includes CO<sub>2</sub> fluxes from deforestation, afforestation, logging and forest degradation (including harvest activity), shifting cultivation (cycle of cutting forest for agriculture, then abandoning), regrowth of forests (following wood harvest or agriculture abandonment), peat burning, and peat drainage. Four bookkeeping approaches (updated estimates each of BLUE (Hansis et al., 2015), OSCAR (Gasser et al., 2020), and H&C2023 (Houghton and Castanho, 2023), and new estimates of LUCE (Qin et al. 2024) were used to quantify gross emissions and gross removals and the resulting net E<sub>LUC</sub>. Emissions from peat burning and peat drainage are added from external datasets, peat drainage being averaged from three spatially explicit independent datasets (see Supplement S.2.1). Uncertainty estimates were derived from the Dynamic Global Vegetation Models (DGVMs) ensemble for the time period prior to 1960, and using for the recent decades an uncertainty range of ±0.7 GtC yr<sup>-1</sup>, which is a semi-quantitative measure for annual and decadal emissions and reflects our best value judgement that there is at least 68% chance (±1σ) that the true land-use change emission lies within the given range, for the range of processes considered here.

The GCB E<sub>LUC</sub> estimates follow the CO<sub>2</sub> flux definition of global carbon cycle models and differ from IPCC definitions adopted in National GHG Inventories (NGHGI) for reporting under the UNFCCC. The latter typically include terrestrial fluxes occurring on all land that countries define as managed, following the IPCC managed land proxy approach (Grassi et al., 2018). This partly includes fluxes due to environmental change (e.g. atmospheric CO<sub>2</sub> increase), which are part of S<sub>LAND</sub> in our definition. As a result, global emission estimates are smaller for NGHGI than for the global carbon budget definition (Grassi et al., 2023). The same is the case for the Food Agriculture Organization (FAO) estimates of carbon fluxes on forest land, which include both anthropogenic and natural fluxes on managed land (Tubiello et al., 2021). We translate the GCB and NGHGI definitions to each other, to provide a comparison of the anthropogenic carbon budget as reported in GCB to the official country reporting to the UNFCCC convention. We further compare these estimates with the net atmosphere-to-land flux from atmospheric inversion systems (see Section 2.7), averaged over managed land only.

E<sub>LUC</sub> contains a range of fluxes that are related to Carbon Dioxide Removal (CDR). CDR is defined as the set of anthropogenic activities that remove CO<sub>2</sub> from the atmosphere, additional to the Earth's natural processes, and store it in durable form, such as in forest biomass and soils, long-lived products, or in geological or ocean

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reservoirs. [Here, we](#) quantify vegetation-based CDR that is implicitly or explicitly captured by land-use fluxes (CDR not based on vegetation is discussed in Section 2.3; IPCC, 2023). We quantify re/afforestation from the [four](#) bookkeeping estimates by separating forest regrowth in shifting cultivation cycles from permanent increases in forest cover (see Supplement [S.2.1](#)). The latter count as CDR, but it should be noted that the permanence of the storage under climate risks such as fire is increasingly questioned. Other CDR activities contained in *E<sub>LUC</sub>* include the transfer of carbon to harvested wood products (HWP), [bioenergy with carbon capture and storage \(BECCS\); and biochar production](#). Note that the [different](#) bookkeeping models represent [HWP](#) with varying details concerning product usage and their lifetimes. Bookkeeping and TRENDY models currently only represent BECCS and biochar with regard to the CO<sub>2</sub> removal through photosynthesis, but [do not account](#) for the durable storage. HWP, BECCS, and biochar are typically counted as CDR when the transfer to the durable storage site occurs and not when the CO<sub>2</sub> is removed from the atmosphere, which complicates a direct comparison to the [GCB](#) approach to quantify annual fluxes to and from the atmosphere. Estimates for CDR through HWP, BECCS, and biochar are thus not indicated in this budget, but can be found elsewhere (see Section 3.2.2).

## 2.2.2 [2024](#) Projection

We project the [2024](#) land-use emissions for BLUE, H&C2023, OSCAR, and LUCE based on their *E<sub>LUC</sub>* estimates for [2023](#) and adding the change in carbon emissions from peat fires and tropical deforestation and degradation fires ([2024](#) emissions relative to [2023](#) emissions) estimated using active fire data (MCD14ML; Giglio et al., 2016). Peat drainage is assumed to be unaltered as it has low interannual variability. More details on the *E<sub>LUC</sub>* methodology can be found in Supplement S.2.

## 2.3 Carbon Dioxide Removal (CDR) not based on vegetation

[While some CDR involves CO<sub>2</sub> fluxes via land-use and is included in \*E<sub>LUC</sub>\*, \(such as afforestation, biochar, HWP, and BECCS\) other CDR occurs through fluxes of CO<sub>2</sub> directly from the air to the geosphere. The majority of this derives from enhanced weathering through the application of crushed rock to soils, with a smaller contribution from Direct Air Carbon Capture and Storage \(DACCS\). We use data from the State of CDR Report \(Smith et al., 2024\), which compiles and harmonises reported removal rates from a combination of existing databases, surveys and novel research. Currently there are no internationally agreed methods for reporting these types of CDR, meaning estimates are based on self-disclosure by projects following their own protocols. As such, the fractional uncertainty on these numbers should be viewed as substantial, and they are liable to change in future years as protocols are harmonised and improved.](#)

## 2.4 Growth rate in atmospheric CO<sub>2</sub> concentration (G<sub>ATM</sub>)

### 2.4.1 [Historical period 1850-2023](#)

The rate of growth of the atmospheric CO<sub>2</sub> concentration is provided for years 1959-[2023](#) by the US National Oceanic and Atmospheric Administration Global Monitoring Laboratory (NOAA/GML; Lan et al., [2024](#)),

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Deleted: ; bioenergy with carbon capture and storage (BECCS); and biochar production.

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1317 which includes recent revisions to the calibration scale of atmospheric CO<sub>2</sub> measurements (WMO-CO<sub>2</sub>-X2019;  
1318 Hall et al., 2021). For the 1959-1979 period, the global growth rate is based on measurements of atmospheric  
1319 CO<sub>2</sub> concentration averaged from the Mauna Loa and South Pole stations, as observed by the CO<sub>2</sub> Program at  
1320 Scripps Institution of Oceanography (Keeling et al., 1976). For the 1980-2021 time period, the global growth  
1321 rate is based on the average of multiple stations selected from the marine boundary layer sites with well-mixed  
1322 background air (Ballantyne et al., 2012), after fitting a smooth curve through the data for each station as a  
1323 function of time, and averaging by latitude band (Masarie and Tans, 1995). The annual growth rate is estimated  
1324 by Lan et al. (2024) from atmospheric CO<sub>2</sub> concentration by taking the average of the most recent December-  
1325 January months corrected for the average seasonal cycle and subtracting this same average one year earlier. The  
1326 growth rate in units of ppm yr<sup>-1</sup> is converted to units of GtC yr<sup>-1</sup> by multiplying by a factor of 2.124 GtC per  
1327 ppm, assuming instantaneous mixing of CO<sub>2</sub> throughout the atmosphere (Ballantyne et al., 2012; Table 1).

1328 The uncertainty around the atmospheric growth rate is due to four main factors. First, the long-term  
1329 reproducibility of reference gas standards (around 0.03 ppm for 1σ from the 1980s; Lan et al., 2024). Second,  
1330 small unexplained systematic analytical errors that may have a duration of several months to two years come  
1331 and go. They have been simulated by randomising both the duration and the magnitude (determined from the  
1332 existing evidence) in a Monte Carlo procedure. Third, the network composition of the marine boundary layer  
1333 with some sites coming or going, gaps in the time series at each site, etc (Lan et al., 2024). The latter uncertainty  
1334 was estimated by NOAA/GML with a Monte Carlo method by constructing 100 "alternative" networks (Masarie  
1335 and Tans, 1995; NOAA/GML, 2019). The second and third uncertainties, summed in quadrature, add up to  
1336 0.085 ppm on average (Lan et al., 2024). Fourth, the uncertainty associated with using the average CO<sub>2</sub>  
1337 concentration from a surface network to approximate the true atmospheric average CO<sub>2</sub> concentration (mass-  
1338 weighted, in 3 dimensions) as needed to assess the total atmospheric CO<sub>2</sub> burden. In reality, CO<sub>2</sub> variations  
1339 measured at the stations will not exactly track changes in total atmospheric burden, with offsets in magnitude  
1340 and phasing due to vertical and horizontal mixing. This effect must be very small on decadal and longer time  
1341 scales, when the atmosphere can be considered well mixed. The CO<sub>2</sub> increase in the stratosphere lags the  
1342 increase (meaning lower concentrations) that we observe in the marine boundary layer, while the continental  
1343 boundary layer (where most of the emissions take place) leads the marine boundary layer with higher  
1344 concentrations. These effects nearly cancel each other. In addition, the growth rate is nearly the same  
1345 everywhere (Ballantyne et al., 2012). We therefore maintain an uncertainty around the annual growth rate based  
1346 on the multiple stations dataset ranges between 0.11 and 0.72 GtC yr<sup>-1</sup>, with a mean of 0.61 GtC yr<sup>-1</sup> for 1959-  
1347 1979 and 0.17 GtC yr<sup>-1</sup> for 1980-2023, when a larger set of stations were available as provided by Lan et al.  
1348 (2024). We estimate the uncertainty of the decadal averaged growth rate after 1980 at 0.02 GtC yr<sup>-1</sup> based on the  
1349 calibration and the annual growth rate uncertainty but stretched over a 10-year interval. For years prior to 1980,  
1350 we estimate the decadal averaged uncertainty to be 0.07 GtC yr<sup>-1</sup> based on a factor proportional to the annual  
1351 uncertainty prior and after 1980 (0.02 \* [0.61/0.17] GtC yr<sup>-1</sup>).

1352 We assign a high confidence to the annual estimates of G<sub>ATM</sub> because they are based on direct measurements  
1353 from multiple and consistent instruments and stations distributed around the world (Ballantyne et al., 2012; Hall  
1354 et al., 2021).

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To estimate the total carbon accumulated in the atmosphere since 1750 or 1850, we use an atmospheric CO<sub>2</sub> concentration of  $278.3 \pm 3$  ppm or  $285.1 \pm 3$  ppm, respectively (Gulev et al., 2021). For the construction of the cumulative budget shown in Figure 3, we use the fitted estimates of CO<sub>2</sub> concentration from Joos and Spahni (2008) to estimate the annual atmospheric growth rate using the conversion factors shown in Table 1. The uncertainty of  $\pm 3$  ppm (converted to  $\pm 1\sigma$ ) is taken directly from the IPCC's AR5 assessment (Ciais et al., 2013). Typical uncertainties in the growth rate in atmospheric CO<sub>2</sub> concentration from ice core data are equivalent to  $\pm 0.1$ - $0.15$  GtC yr<sup>-1</sup> as evaluated from the Law Dome data (Etheridge et al., 1996) for individual 20-year intervals over the period from 1850 to 1960 (Bruno and Joos, 1997).

#### 2.4.2 2024 projection

We provide an assessment of G<sub>ATM</sub> for 2024 as the average of two methods. The GCB regression method models monthly global-average atmospheric CO<sub>2</sub> concentrations and derives the increment and annual average from these. The model uses lagged observations of concentration (Lan et al., 2024): both a 12-month lag, and the lowest lag that will allow model prediction to produce an estimate for the following January, recalling that the G<sub>ATM</sub> increment is derived from December/January pairs. The largest driver of interannual changes is the ENSO signal (Betts et al., 2016), so the monthly ENSO 3.4 index (Huang et al., 2023) is included in the model. Given the natural lag between sea-surface temperatures and effects on the biosphere, and in turn effects on globally mixed atmospheric CO<sub>2</sub> concentration, a lagged ENSO index is used, and we use both a 5-month and a 6-month lag. The combination of the two lagged ENSO values helps reduce possible effects of noise in a single month. To help characterise the seasonal variation, we add month as a categorical variable. Finally, we flag the period affected by the Pinatubo eruption (August 1991 - November 1993) as a categorical variable. Note that while emissions of CO<sub>2</sub> are the largest driver of the trend in atmospheric CO<sub>2</sub> concentration, our goal here is to predict divergence from that trend. Because changes in emissions from year to year are relatively minor in comparison to total emissions, this has little effect on the variation of concentration from the trend line. Even the relatively large drop in emissions in 2020 due to the COVID-19 pandemic does not cause any problems for the model.

We also use the multi-model mean and uncertainty of the 2024 G<sub>ATM</sub> estimated by the ESMs prediction system (see Section 2.9). We then take the average of the GCB regression and ESMs G<sub>ATM</sub> estimates, with their respective uncertainty combined quadratically.

Similarly, the projection of the 2024 global average CO<sub>2</sub> concentration (in ppm), is calculated as the average of the estimates from the two methods. For the GCB regression method, it is the annual average of global concentration over the 12 months of 2024; for the ESMs, it is the observed global average CO<sub>2</sub> concentration for 2023 plus the annual increase in 2024 of the global average CO<sub>2</sub> concentration predicted by the ESMs multi-model mean.

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1419 **2.5 Ocean CO<sub>2</sub> sink**

1420 **2.5.1 Historical period 1850-2023**

1421 The reported estimate of the global ocean anthropogenic CO<sub>2</sub> sink  $S_{\text{OCEAN}}$  is derived as the average of two  
1422 estimates. The first estimate is derived as the mean over an ensemble of ten global ocean biogeochemistry  
1423 models (GOBMs, Table 4 and Table S2). The second estimate is obtained as the mean over an ensemble of ~~eight~~  
1424 surface ocean  $f\text{CO}_2$ -observation-based data-products (Table 4 and Table S3). ~~A ninth  $f\text{CO}_2$ -product (JExP-FFN-~~  
1425 ~~U) is shown, but is not included in the ensemble average as it differs from the other products by adjusting the~~  
1426 ~~flux to a cool, salty ocean surface skin. In previous editions of the GCB, this product was following the Watson~~  
1427 ~~et al. (2020) method but has been updated following the method of Dong et al. (2022, see Supplement S.3.1 for~~  
1428 ~~a discussion). The GOBMs simulate both the natural and anthropogenic CO<sub>2</sub> cycles in the ocean. They constrain~~  
1429 ~~the anthropogenic air-sea CO<sub>2</sub> flux (the dominant component of  $S_{\text{OCEAN}}$ ) by the transport of carbon into the~~  
1430 ~~ocean interior, which is also the controlling factor of present-day ocean carbon uptake in the real world. They~~  
1431 ~~cover the full globe and all seasons and were evaluated against surface ocean carbon observations, suggesting~~  
1432 ~~they are suitable to estimate the annual ocean carbon sink (Hauck et al., 2020). The  $f\text{CO}_2$ -products are tightly~~  
1433 ~~linked to observations of  $f\text{CO}_2$  (fugacity of CO<sub>2</sub>, which equals  $p\text{CO}_2$  corrected for the non-ideal behaviour of the~~  
1434 ~~gas; Pfeil et al., 2013), which carry imprints of temporal and spatial variability, but are also sensitive to~~  
1435 ~~uncertainties in gas-exchange parameterizations and data-sparsity (Fay et al., 2021, Gloege et al., 2021, Hauck~~  
1436 ~~et al., 2023a). Their asset is the assessment of the mean spatial pattern of variability and its seasonality (Hauck~~  
1437 ~~et al., 2020, Gloege et al. 2021, Hauck et al., 2023a). To benchmark trends derived from the  $f\text{CO}_2$ -products, we~~  
1438 ~~additionally performed a model subsampling exercise following Hauck et al. (2023a, see section S3). In~~  
1439 ~~addition, two diagnostic ocean models are used to estimate  $S_{\text{OCEAN}}$  over the industrial era (1781-1958).~~

1440 The global  $f\text{CO}_2$ -based flux estimates were adjusted to remove the pre-industrial ocean source of CO<sub>2</sub> to the  
1441 atmosphere of  $0.65 \pm 0.3 \text{ GtC yr}^{-1}$  from river input to the ocean (Regnier et al., 2022), to satisfy our definition of  
1442  $S_{\text{OCEAN}}$  (Hauck et al., 2020). The river flux adjustment was distributed over the latitudinal bands using the  
1443 regional distribution of Lacroix et al. (2020; North:  $0.14 \text{ GtC yr}^{-1}$ , Tropics:  $0.42 \text{ GtC yr}^{-1}$ , South:  $0.09 \text{ GtC yr}^{-1}$ ).  
1444 Acknowledging that this distribution is based on only one model, the advantage is that a gridded field is  
1445 available, and the river flux adjustment can be calculated for the three latitudinal bands and the RECCAP  
1446 regions (REgional Carbon Cycle Assessment and Processes (RECCAP2; Ciais et al., 2020, Poulter et al., 2022,  
1447 DeVries et al., 2023). This ~~dataset~~ suggests that more of the riverine outgassing is located in the tropics than in  
1448 the Southern Ocean, and is thus opposed to the previously used ~~dataset~~ of Aumont et al. (2001). Accordingly, the  
1449 regional distribution is associated with a major uncertainty in addition to the large uncertainty around the global  
1450 estimate (Crisp et al., 2022; Gruber et al., 2023). Anthropogenic perturbations of river carbon and nutrient  
1451 transport to the ocean are not considered (see Section 2.10 and Supplement S.6.3).

1452 We derive  $S_{\text{OCEAN}}$  from GOBMs by using a simulation (sim A) with historical forcing of climate and  
1453 atmospheric CO<sub>2</sub> ~~from GCB (Section 2.4)~~, accounting for model biases and drift from a control simulation (sim  
1454 B) with constant atmospheric CO<sub>2</sub> and normal year climate forcing. A third simulation (sim C) with historical

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1469 atmospheric CO<sub>2</sub> increase and normal year climate forcing is used to attribute the ocean sink to CO<sub>2</sub> (sim C  
1470 minus sim B) and climate (sim A minus sim C) effects. A fourth simulation (sim D; historical climate forcing  
1471 and constant atmospheric CO<sub>2</sub>) is used to compare the change in anthropogenic carbon inventory in the interior  
1472 ocean (sim A minus sim D) to the observational estimate of Gruber et al. (2019) with the same flux components  
1473 (steady state and non-steady state anthropogenic carbon flux). The  $f\text{CO}_2$ -products are adjusted with respect to  
1474 their original publications to represent the full ice-free ocean area, including coastal zones and marginal seas,  
1475 when the area coverage is below 99%. This is done by either area filling following Fay et al. (2021) or a simple  
1476 scaling approach. GOBMs and  $f\text{CO}_2$ -products fall within the observational constraints over the 1990s ( $2.2 \pm 0.7$   
1477 GtC yr<sup>-1</sup>, Ciais et al., 2013) before and after applying adjustments.

1478  $\text{SOCEAN}$  is calculated as the average of the GOBM ensemble mean and the  $f\text{CO}_2$ -product ensemble mean from  
1479 1990 onwards. Prior to 1990, it is calculated as the GOBM ensemble mean plus half of the offset between  
1480 GOBMs and  $f\text{CO}_2$ -products ensemble means over 1990-2001.

1481 We assign an uncertainty of  $\pm 0.4$  GtC yr<sup>-1</sup> to the ocean sink based on a combination of random (ensemble  
1482 standard deviation) and systematic uncertainties (GOBMs bias in anthropogenic carbon accumulation,  
1483 previously reported uncertainties in  $f\text{CO}_2$ -products; see Supplement S.3.4). While this approach is consistent  
1484 within the GCB, an independent uncertainty assessment of the  $f\text{CO}_2$ -products alone suggests a somewhat larger  
1485 uncertainty of up to 0.7 GtC yr<sup>-1</sup> (Ford et al. 2024, accepted). We assess a medium confidence level to the  
1486 annual ocean CO<sub>2</sub> sink and its uncertainty because it is based on multiple lines of evidence, it is consistent with  
1487 ocean interior carbon estimates (Gruber et al., 2019, see Section 3.6.5) and the interannual variability in the  
1488 GOBMs and data-based estimates is largely consistent and can be explained by climate variability. We refrain  
1489 from assigning a high confidence because of the deviation between the GOBM and  $f\text{CO}_2$ -product trends  
1490 between around 2002 and 2020. More details on the  $\text{SOCEAN}$  methodology can be found in Supplement S.3.

1491 **2.5.2 2024 Projection**

1492 The ocean CO<sub>2</sub> sink forecast for the year 2024 is based on the annual historical time-series and our estimated  
1493 2024 atmospheric CO<sub>2</sub> concentration (Lan et al 2024), the historical and our estimated 2024 annual global fossil  
1494 fuel emissions from this year's carbon budget, and the spring (March, April, May) Oceanic Niño Index (ONI)  
1495 (NCEP, 2024). Using a non-linear regression approach, i.e., a feed-forward neural network, atmospheric CO<sub>2</sub>,  
1496 ONI, and the fossil fuel emissions are used as training data to best match the annual ocean CO<sub>2</sub> sink (i.e.  
1497 combined  $\text{SOCEAN}$  estimate from GOBMs and data products) from 1959 through 2023 from this year's carbon  
1498 budget. Using this relationship, the 2024  $\text{SOCEAN}$  can then be estimated from the projected 2024 input data using  
1499 the non-linear relationship established during the network training. To avoid overfitting, the neural network was  
1500 trained with a variable number of hidden neurons (varying between 2-5) and 20% of the randomly selected  
1501 training data were withheld for independent internal testing. Based on the best output performance (tested using  
1502 the 20% withheld input data), the best performing number of neurons was selected. In a second step, we trained  
1503 the network 10 times using the best number of neurons identified in step 1 and different sets of randomly  
1504 selected training data. The mean of the 10 trainings is considered our best forecast, whereas the standard

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deviation of the 10 ensembles provides a first order estimate of the forecast uncertainty. This uncertainty is then combined with the  $S_{OCEAN}$  uncertainty ( $0.4 \text{ GtC yr}^{-1}$ ) to estimate the overall uncertainty of the 2024 projection. As an additional line of evidence, we also assess the 2024 atmosphere-ocean carbon flux from the ESM prediction system (see Section 2.9).

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## 2.6 Land CO<sub>2</sub> sink

### 2.6.1 Historical Period 1850-2023

The terrestrial land sink ( $S_{LAND}$ ) is thought to be due to the combined effects of rising atmospheric CO<sub>2</sub>, increasing N inputs, and climate change, on plant growth, and terrestrial carbon storage.  $S_{LAND}$  does not include land sinks directly resulting from land-use and land-use change (e.g., regrowth of vegetation) as these are part of the land-use flux ( $E_{LUC}$ ), although system boundaries make it difficult to attribute exactly CO<sub>2</sub> fluxes on land between  $S_{LAND}$  and  $E_{LUC}$  (Erb et al., 2013).

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$S_{LAND}$  is estimated from the multi-model mean of 20 DGVMs (Table 4 and Table S1). DGVMs simulations include all climate variability and CO<sub>2</sub> effects over land. In addition to the carbon cycle represented in all DGVMs, 14 models also account for the nitrogen cycle and hence can include the effect of N inputs on  $S_{LAND}$ .

Deleted: S1) with an additional comparison of DGVMs with a data-driven, carbon data model framework (CARDAMOM) (Bloom and Williams, 2015; Bloom et al., 2016), see Supplement S4....

The DGVMs estimate of  $S_{LAND}$  does not include the export of carbon to aquatic systems or its historical perturbation, which is discussed in Supplement S.6.3. DGVMs need to meet several criteria to be included in this assessment. In addition, we use the International Land Model Benchmarking system (ILAMB; Collier et al., 2018) for the DGVMs evaluation (see Supplement S.4.2), with an additional comparison of DGVMs with a data-informed, Bayesian model-data fusion framework (CARDAMOM) (Bloom and Williams, 2015; Bloom et al., 2016). The uncertainty on  $S_{LAND}$  is taken from the DGVMs standard deviation. More details on the  $S_{LAND}$  methodology can be found in Supplement S.4.

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### 2.6.2 2024 Projection

Like for the ocean forecast, the land CO<sub>2</sub> sink ( $S_{LAND}$ ) forecast for the year 2024 is based on the annual historical (Lan et al., 2024) and our estimated 2024 atmospheric CO<sub>2</sub> concentration, historical and our estimated 2024 annual global fossil fuel emissions from this year's carbon budget, and the summer (June, July, August) ONI (NCEP, 2024). All training data are again used to best match  $S_{LAND}$  from 1959 through 2023 from this year's carbon budget using a feed-forward neural network. To avoid overfitting, the neural network was trained with a variable number of hidden neurons (varying between 2-15), larger than for  $S_{OCEAN}$  prediction due to the stronger land carbon interannual variability. As done for  $S_{OCEAN}$ , a pre-training selects the optimal number of hidden neurons based on 20% withheld input data, and in a second step, an ensemble of 10 forecasts is produced to provide the mean forecast plus uncertainty. This uncertainty is then combined with the  $S_{LAND}$  uncertainty for 2023 ( $1.0 \text{ GtC yr}^{-1}$ ) to estimate the overall uncertainty of the 2024 projection.

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573 2.7 Atmospheric inversion estimate

1574 The world-wide network of in-situ atmospheric measurements and satellite derived atmospheric CO<sub>2</sub> column  
1575 (xCO<sub>2</sub>) observations put a strong constraint on changes in the atmospheric abundance of CO<sub>2</sub>. This is true  
1576 globally (hence our large confidence in G<sub>ATM</sub>), but also in regions with sufficient observational density found  
1577 mostly in the extra-tropics. This allows atmospheric inversion methods to constrain the magnitude and location  
1578 of the combined total surface CO<sub>2</sub> fluxes from all sources, including fossil and land-use change emissions and  
1579 land and ocean CO<sub>2</sub> fluxes. The inversions assume E<sub>FOS</sub> to be well known, and they solve for the spatial and  
1580 temporal distribution of land and ocean fluxes from the residual gradients of CO<sub>2</sub> between stations that are not  
1581 explained by fossil fuel emissions. By design, such systems thus close the carbon balance (B<sub>IM</sub> = 0) and thus  
1582 provide an additional perspective on the independent estimates of the ocean and land fluxes.

1583 This year's release includes fourteen inversion systems that are described in Table S4. Each system is rooted in  
1584 Bayesian inversion principles but uses different methodologies. These differences concern the selection of  
1585 atmospheric CO<sub>2</sub> data or xCO<sub>2</sub>, and the choice of a-priori fluxes to refine. They also differ in spatial and  
1586 temporal resolution, assumed correlation structures, and mathematical approach of the models (see references in  
1587 Table S4 for details). Importantly, the systems use a variety of transport models, which was demonstrated to be  
1588 a driving factor behind differences in atmospheric inversion-based flux estimates, and specifically their  
1589 distribution across latitudinal bands (Gaubert et al., 2019; Schuh et al., 2019). [Eight inversion systems used](#)  
1590 [surface observations from the global measurement network \(Schuldt et al., 2023, 2024\)](#). Six inversion systems  
1591 (CAMS-[FT24r1](#), CMS-flux, GONGGA, [COLA](#), GCASv2, [NTFVAR](#)) used satellite xCO<sub>2</sub> retrievals from  
1592 GOSAT and/or OCO-2, scaled to the WMO 2019 calibration scale, [of which three](#) inversions this year (CMS-  
1593 Flux, [COLA](#), [NTFVAR](#)) used these xCO<sub>2</sub> datasets in addition to the in-situ observational CO<sub>2</sub> mole fraction  
1594 records.

1595 The original products delivered by the inverse modellers were modified to facilitate the comparison to the other  
1596 elements of the budget, specifically on two accounts: (1) global total fossil fuel emissions including cement  
1597 carbonation CO<sub>2</sub> uptake, and (2) riverine CO<sub>2</sub> transport. [We note that with these adjustments the inverse results](#)  
1598 no longer represent the net atmosphere-surface exchange over land/ocean areas as sensed by atmospheric  
1599 observations. Instead, for land, they become the net uptake of CO<sub>2</sub> by vegetation and soils that is not exported  
1600 by fluvial systems, similar to the DGVMs estimates. For oceans, they become the net uptake of anthropogenic  
1601 CO<sub>2</sub>, similar to the GOBMs estimates.

1602 The inversion systems prescribe global fossil fuel emissions based on e.g. the GCP's Gridded Fossil Emissions  
1603 Dataset versions [2024.0](#) (GCP-GridFED; Jones et al., [2024a](#)), which are updates to GCP-GridFEDv2021  
1604 presented by Jones et al. (2021b). GCP-[GridFEDv2024.0](#) scales gridded estimates of CO<sub>2</sub> emissions from  
1605 EDGARv4.3.2 (Janssens-Maenhout et al., 2019) within national territories to match national emissions  
1606 estimates provided by the GCB for the years 1959-[2023](#), which were compiled following the methodology  
1607 described in Section 2.1. Small differences between the systems due to for instance regridding to the transport  
1608 model resolution, or use of different fossil fuel emissions [than GCP-GridFEDv2024.0](#), are adjusted in the

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inverse estimates presented in the text and figures.

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latitudinal partitioning we present, to ensure agreement with the estimate of  $E_{FOS}$  in this budget. We also note that the ocean fluxes used as prior by 8 out of 14 inversions are part of the suite of the ocean process model or  $fCO_2$ -products listed in Section 2.5. Although these fluxes are further adjusted by the atmospheric inversions, (except for Jena CarboScope), it makes the inversion estimates of the ocean fluxes not completely independent of  $SO_{CEAN}$  assessed here.

To facilitate comparisons to the independent  $SO_{CEAN}$  and  $S_{LAND}$ , we used the same adjustments for transport and outgassing of carbon transported from land to ocean, as done for the observation-based estimates of  $SO_{CEAN}$  (see Supplement S.3).

The atmospheric inversions are evaluated using vertical profiles of atmospheric  $CO_2$  concentrations (Figure S5). More than 30 aircraft programs over the globe, either regular programs or repeated surveys over at least 9 months (except for SH programs), have been used to assess system performance (with space-time observational coverage sparse in the SH and tropics, and denser in NH mid-latitudes; Table S8). The fourteen systems are compared to the independent aircraft  $CO_2$  measurements between 2 and 7 km above sea level between 2001 and 2023. Results are shown in Figure S5 and discussed in Supplement S.5.2.

With a relatively small ensemble of systems that cover at least one full decade ( $N=10$ ), and which moreover share some a-priori fluxes used with one another, or with the process-based models, it is difficult to justify using their mean and standard deviation as a metric for uncertainty across the ensemble. We therefore report their full range (min-max) without their mean. More details on the atmospheric inversion methodology can be found in Supplement S.5.

## 2.8 Atmospheric oxygen based estimate

Long-term atmospheric  $O_2$  and  $CO_2$  observations allow estimation of the global ocean and land carbon sinks, due to the coupling of  $O_2$  and  $CO_2$  with distinct exchange ratios for fossil fuel emissions and land uptake, and uncoupled  $O_2$  and  $CO_2$  ocean exchange (Keeling and Manning, 2014). The global ocean and net land carbon sinks were calculated following methods and constants used in Keeling and Manning (2014) but modified to also include the effective  $O_2$  source from metal refining (Battle et al., 2023). For the exchange ratio of the net land sink at value of 1.05 is used, following Resplandy et al. (2019). For fossil fuels, the following values are used: gas: 1.95 (+/-) 0.04, liquid: 1.44, (+/-) 0.03, solid: 1.17 (+/-) 0.03, cement: 0 (+/-) 0, gas flaring: 1.98 (+/-) 0.07 (Keeling, 1988). Atmospheric  $O_2$  is observed as  $\delta(O_2/N_2)$  and combined with  $CO_2$  mole fraction observations into Atmospheric Potential Oxygen (APO, Stephens et al., 1998). The APO observations from 1990 to 2024 were taken from a weighted average of flask records from three stations in the Scripps  $O_2$  program network (Alert, Canada (ALT), La Jolla, California (LJO), and Cape Grim, Australia (CGO), weighted per Keeling and Manning (2014). Observed  $CO_2$  was taken from the globally averaged marine surface annual mean growth rate from the NOAA/GML Global Greenhouse Gas Reference Network (Lan et al., 2024). The  $O_2$  source from ocean warming is based on ocean heat content from updated data from NOAA/NCEI (Levitus et al., 2012). The effective  $O_2$  source from metal refining is based on production data from Bray (2020), Flanagan (2021), and Tuck (2022). Uncertainty was determined through a Monte Carlo approach with 20,000 iterations, using

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1673 uncertainties prescribed in Keeling and Manning (2014), including observational uncertainties from Keeling et  
1674 al. (2007) and autoregressive errors in fossil fuel emissions (Ballantyne et al., 2015). The reported uncertainty is  
1675 1 standard deviation of the ensemble. [The difference between the atmospheric O<sub>2</sub> estimate for GCB2023 is due](#)  
1676 [to a revision to the Scripps O<sub>2</sub> program CO<sub>2</sub> data. As for the atmospheric inversions, the O<sub>2</sub> based estimates also](#)  
1677 [closes the carbon balance \( \$B\_M = 0\$ \) by design and provides another independent estimate of the ocean and land](#)  
1678 [fluxes. Note that the O<sub>2</sub> method requires a correction for global air-sea O<sub>2</sub> flux, which has the largest uncertainty](#)  
1679 [at annual time scales, but which is still non negligible for decadal estimates \(Nevison et al., 2008\).](#)

1680 **2.9 Earth System Models estimate**

1681 Reconstructions and predictions from decadal prediction systems based on Earth system models (ESMs) provide  
1682 a novel line of evidence in assessing the atmosphere-land and atmosphere-ocean carbon fluxes in the past  
1683 decades and predicting their changes for the current year. The decadal prediction systems based on ESMs used  
1684 here consist of three sets of simulations: (i) uninitialized freely evolving historical simulations (1850-2014); (ii)  
1685 assimilation reconstruction incorporating observational data into the model ([1960-2023](#)); (iii) [initialised](#)  
1686 prediction simulations for the 1981-[2024](#) period, starting every year from initial states obtained from the above  
1687 assimilation simulations. The assimilations are designed to reconstruct the actual evolution of the Earth system  
1688 by assimilating essential fields from data products. The assimilations' states, which are expected to be close to  
1689 observations, are used to start the [initialised](#) prediction simulations used for the current year ([2024](#)) global  
1690 carbon budget. Similar [initialised](#) prediction simulations starting every year (Nov. 1st or Jan. 1st) over the 1981-  
1691 [2023](#) period (i.e., hindcasts) are also performed for predictive skill quantification and for bias correction. More  
1692 details on the illustration of a decadal prediction system based on an ESM can refer to Figure 1 of Li et al.  
1693 (2023).

1694 By assimilating physical atmospheric and oceanic data products into the ESMs, the models are able to reproduce  
1695 the historical variations of the atmosphere-sea CO<sub>2</sub> fluxes, atmosphere-land CO<sub>2</sub> fluxes, and atmospheric CO<sub>2</sub>  
1696 growth rate (Li et al., 2016, 2019; Lovenduski et al., 2019a,b; Ilyina et al., 2021; Li et al., 2023). Furthermore,  
1697 the ESM-based predictions have proven their skill in predicting the air-sea CO<sub>2</sub> fluxes for up to 6 years, the air-  
1698 land CO<sub>2</sub> fluxes and atmospheric CO<sub>2</sub> growth for 2 years (Lovenduski et al., 2019a,b; Ilyina et al., 2021; Li et  
1699 al., 2023). The reconstructions from the fully coupled model simulations ensure a closed budget within the Earth  
1700 system, i.e., no budget imbalance term.

1701 [Five](#) ESMs, i.e., CanESM5 (Swart et al., 2019; Sospedra-Alfonso et al., 2021), [EC-Earth3-CC](#) (Döscher et al.  
1702 [2021; Bilbao et al., 2021; Bernardello et al., 2024](#)), IPSL-CM6A-CO2-LR (Boucher et al., 2020), MIROC-ES2L  
1703 (Watanabe et al., 2020), and MPI-ESM1-2-LR (Mauritsen et al., 2019; Li et al., 2023), have performed the set of  
1704 prediction simulations. Each ESM uses a different assimilation method and combination of data products  
1705 incorporated in the system, more details on the models configuration can be found in [Table 4](#) and [Supplementary](#)  
1706 [Table S5](#). The ESMs use external forcings from the Coupled Model Intercomparison Project Phase 6 (CMIP6)  
1707 historical ([1960-2014](#)) plus SSP2-4.5 baseline and CovidMIP two-year blip scenario (2015-[2024](#)) (Eyring et al.,  
1708 2016; [Lamboll et al., 2021](#)). The CO<sub>2</sub> emissions forcing from 2015-[2024](#) are substituted by GCB-GridFED

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1724 [\(2024.0, Jones et al., 2024a\)](#) to provide a [consistent CO<sub>2</sub>](#) forcing. Reconstructions of atmosphere-ocean CO<sub>2</sub>  
1725 fluxes (SOCEAN) and atmosphere-land CO<sub>2</sub> fluxes (SLAND-ELUC) for the time period from [1960-2023](#) are assessed  
1726 here. Predictions of the atmosphere-ocean CO<sub>2</sub> flux, atmosphere-land CO<sub>2</sub> flux, and atmospheric CO<sub>2</sub> growth for  
1727 [2024](#) are calculated based on the predictions at a lead time of 1 year. The predictions are [bias-corrected](#) using the  
1728 1985-2014 climatology mean of GCB2022 (Friedlingstein et al., 2022), more details on methods can be found in  
1729 Boer et al. (2016) and Li et al. (2023). The ensemble size of initialized prediction simulations is 10, and the  
1730 ensemble mean for each individual model is used here. The ESMs are used here to support the assessment of  
1731 SOCEAN and net atmosphere-land CO<sub>2</sub> flux (SLAND - ELUC) over the [1960-2023](#) period, and to provide an estimate  
1732 of the [2024](#) projection of GATM.

### 1733 2.10 Processes not included in the global carbon budget

1734 The contribution of anthropogenic CO and CH<sub>4</sub> to the global carbon budget is not fully accounted for in Eq. (1)  
1735 and is described in Supplement S.6.1. The contributions to CO<sub>2</sub> emissions of decomposition of carbonates not  
1736 accounted for is described in Supplement S.6.2. The contribution of anthropogenic changes in river fluxes is  
1737 conceptually included in Eq. (1) in SOCEAN and in SLAND, but it is not represented in the process models used to  
1738 quantify these fluxes. This effect is discussed in Supplement S.6.3. Similarly, the loss of additional sink capacity  
1739 from reduced forest cover is missing in the combination of approaches used here to estimate both land fluxes  
1740 (ELUC and SLAND) and its potential effect is discussed and quantified in Supplement S.6.4.

## 1741 3 Results

1742 For each component of the global carbon budget, we present results for three different time periods: the full  
1743 historical period, from 1850 to [2023](#), the decades in which we have atmospheric concentration records from  
1744 Mauna Loa (1960-[2023](#)), a specific focus on last year ([2023](#)), and the projection for the current year ([2024](#)).  
1745 Subsequently, we assess the estimates of the budget components of the last decades against the top-down  
1746 constraints from inverse modelling of atmospheric observations, the land/ocean partitioning derived from the  
1747 atmospheric O<sub>2</sub> measurements, and the budget components estimates from the ESMs assimilation simulations.  
1748 Atmospheric inversions further allow for an assessment of the budget components with a regional breakdown of  
1749 land and ocean sinks.

### 1750 3.1 Fossil CO<sub>2</sub> Emissions

#### 1751 3.1.1 Historical period 1850-[2023](#)

1752 Cumulative fossil CO<sub>2</sub> emissions for 1850-[2023](#) were [490 ± 25 GtC](#), including the cement carbonation sink  
1753 (Figure 3, Table 8, with all cumulative numbers rounded to the nearest 5GtC). In this period, 46% of global  
1754 fossil CO<sub>2</sub> emissions came from coal, 35% from oil, 15% from natural gas, 3% from decomposition of  
1755 carbonates, and 1% from flaring. In 1850, the UK stood for 62% of global fossil CO<sub>2</sub> emissions. In [1893](#) the  
1756 combined cumulative emissions of the current members of the European Union reached and subsequently  
1757 surpassed the level of the UK. Since 1917 US cumulative emissions have been the largest. Over the entire

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774	period 1850-2023, US cumulative emissions amounted to 120GtC (24% of world total), the EU's to 80 GtC	Deleted: 2022...023, US cumulative emissions amounted to 115GtC...20GtC (24% of world total), the EU's to 80 GtC
775	(46%), China's to 75 GtC (15%), and India's to 15 GtC (3%).	(17...6%), China's to 70 ... [33]
776	In addition to the estimates of fossil CO <sub>2</sub> emissions that we provide here (see Section 2.1), there are three global	Deleted: Methods...ection 2.1), there are three global
777	datasets with long time series that include all sources of fossil CO <sub>2</sub> emissions: CDIAC-FF (Hefner and Marland,	datasets with long time series that include all sources of fossil
778	2024), CEDS version 2024 07 08 (Hoesly et al., 2024) and PRIMAP-hist version 2.6 (Gütschow et al., 2016;	CO <sub>2</sub> emissions: CDIAC-FF (Gilfillan...efner and Marland,
779	Gütschow et al., 2024), although these datasets are not entirely independent from each other (Andrew, 2020a).	2021...024), CEDS version v_2021_04_21...024_07_08
780	CEDS has cumulative emissions over 1750-2022 at 480 GtC, CDIAC-FF has 481 GtC, GCP 484 GtC,	(Hoesly et al., 2018; O'Rourke et al., 2021...024) and
781	PRIMAP-hist CR 490 GtC, and PRIMAP-hist TR 492 GtC. CDIAC-FF excludes emissions from lime	PRIMAP-hist version 2.4.2... (Gütschow et al., 2016;
782	production. CEDS estimates higher emissions from international shipping in recent years, while PRIMAP-hist	Gütschow and Prflüger, 2023...t al., 2024), although these
783	has higher fugitive emissions than the other datasets. However, in general these four datasets are in relative	datasets are not entirely independent from each other
784	agreement as to total historical global emissions of fossil CO <sub>2</sub> .	(Andrew, 2020a). CDIAC-FF...EDS has the lowest
		cumulative emissions over 1750-2018...022 at 440...80 GtC,
		GtC, CDIAC-FF has 481 GtC, GCP has 444 GtC, CEDS
		445...84 GtC, PRIMAP-hist TP 453...R 490 GtC, and
		PRIMAP-hist CR 452...R 492 GtC. CDIAC-FF excludes
		emissions from lime production. CEDS has ... [34]
785	3.1.2 Recent period 1960-2023	Deleted: 2022
786	Global fossil CO <sub>2</sub> emissions, E <sub>fos</sub> (including the cement carbonation sink), have increased every decade from an	Formatted: Outline numbered + Level: 3 + Numbering
787	average of 3.0 ± 0.2 GtC yr <sup>-1</sup> for the decade of the 1960s to an average of 9.7 ± 0.5 GtC yr <sup>-1</sup> during 2014-2023,	Style: 1, 2, 3, ... + Start at: 1 + Alignment: Left + Aligned at:
788	(Table 7, Figure 2 and Figure 5). The growth rate in these emissions decreased between the 1960s and the	0 cm + Indent at: 1.27 cm
789	1990s, from 4.3% yr <sup>-1</sup> in the 1960s (1960-1969), 3.2% yr <sup>-1</sup> in the 1970s (1970-1979), 1.6% yr <sup>-1</sup> in the 1980s	Deleted: 6... ± 0.5 GtC yr <sup>-1</sup> during 2013-2022 ... [35]
790	(1980-1989), to 1.0% yr <sup>-1</sup> in the 1990s (1990-1999). After this period, the growth rate began increasing again in	Formatted: Pattern: Clear
791	the 2000s at an average growth rate of 2.8% yr <sup>-1</sup> , decreasing to 0.6% yr <sup>-1</sup> for the last decade (2014-2023).	Deleted: 5...% yr <sup>-1</sup> for the last decade (2013-2022...014-
792	China's emissions increased by +1.9% yr <sup>-1</sup> on average over the last 10 years dominating the global trend, and	2023). China's emissions increased by +1.6...% yr <sup>-1</sup> on
793	India's emissions increased by +3.6% yr <sup>-1</sup> , while emissions decreased in EU27 by 2.1% yr <sup>-1</sup> , and in the USA by	average over the last 10 years dominating the global trend,
794	1.2% yr <sup>-1</sup> . Figure 6 illustrates the spatial distribution of fossil fuel emissions for the 2014-2023 period.	and India's emissions increased by +3.5...% yr <sup>-1</sup> , while
		emissions decreased in EU27 by -...1.7... yr <sup>-1</sup> , and in the
		USA by -...0...% yr <sup>-1</sup> . Figure 6 illustrates the spatial
		distribution of fossil fuel emissions for the 2013-2022 ... [36]
795	E <sub>fos</sub> reported here includes the uptake of CO <sub>2</sub> by cement via carbonation which has increased with increasing	Deleted: 18...0 MtC yr <sup>-1</sup> (0.018...2 GtC yr <sup>-1</sup> ) in the 1960s to
796	stocks of cement products, from an average of 20 MtC yr <sup>-1</sup> (0.02 GtC yr <sup>-1</sup> ) in the 1960s to an average of 200MtC	an average of 197 MtC...00MtC yr <sup>-1</sup> (0.197... GtC yr <sup>-1</sup> )
797	yr <sup>-1</sup> (0.2 GtC yr <sup>-1</sup> ) during 2014-2023 (Figure 5).	during 2013-2022... ... [37]
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798	3.1.3 Final year 2023	Deleted: 2022
799	Global fossil CO <sub>2</sub> emissions were slightly higher, 1.4%, in 2023 than in 2022, with an increase of 0.14 GtC to	Formatted: Outline numbered + Level: 3 + Numbering
800	reach 10.1 ± 0.5 GtC (including the 0.21 GtC cement carbonation sink) in 2023 (Figure 5), distributed among	Style: 1, 2, 3, ... + Start at: 1 + Alignment: Left + Aligned at:
801	coal (41%), oil (32%), natural gas (21%), cement (4%), flaring (<1%), and others (<1%). Compared to 2022, the	0 cm + Indent at: 1.27 cm
802	2023 emissions from coal, oil, and gas increased by 1.4%, 2.5%, and 0.1% respectively, while emissions from	Deleted: 0.9...4%, in 2022...023 than in 2021...022, with an
803	cement decreased by 2%. All annual growth rates presented are adjusted for the leap year, unless stated	increase of less than ...1...4 GtC to reach 9.9...0.1 ± 0.5
804	otherwise.	(including the 0.2...1 GtC cement carbonation sink) in
		2022...023 (Figure 5), distributed among coal (41%), oil
		(32%), natural gas (21%), cement (4%), flaring (...<1%), and
		others (<1%). Compared to 2022, the previous year,
		2022...023 emissions from coal, oil, and oil...as increased by
		1.6%...%, 2.5%, and 3.2...1% respectively, while emissions
		from gas and ...ement respectively ...creased by 2.2% and
		5.1... ... [38]
805	In 2023, the largest absolute contributions to global fossil CO <sub>2</sub> emissions were from China (31%), the USA	Deleted: 2022...023, the largest absolute contributions to
806	(13%), India (8%), and the EU27 (7%). These four regions account for 59% of global fossil CO <sub>2</sub> emissions,	global fossil CO <sub>2</sub> emissions were from China (31%), the USA
807	while the rest of the world contributed 41%, including international aviation and marine bunker fuels (3% of the	(14...3%), India (8%), and the EU27 (7%). These four
		regions account for 59% of global fossil CO <sub>2</sub> emissions,
		while the rest of the world contributed 41%, including
		international aviation and marine bunker fuels (2.6 ... [39]



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total). Growth rates for these countries from 2022 to 2023 were 4.9% (China), 3.3% (USA), 8.4% (EU27), and 8.2% (India), with +0.7% for the rest of the world, including international aviation and marine bunker fuels (+9.5%). The per-capita fossil CO<sub>2</sub> emissions in 2023 were 1.3 tC person<sup>-1</sup> yr<sup>-1</sup> for the globe, and were 3.9 (USA), 2.3 (China), 1.5 (EU27) and 0.6 (India) tC person<sup>-1</sup> yr<sup>-1</sup> for the four highest emitters (Figure 5).

**Deleted:** 2021 to ...022 to 2023 were 0.5...9% (China), 0.5...3.3% (USA), -1.6...4% (EU27), and 5...2% (India), with +0.9...% for the rest of the world... including international aviation and marine bunker fuels (+9.5%). The per-capita fossil CO<sub>2</sub> emissions in 2022...023 were 1.3 tC person<sup>-1</sup> yr<sup>-1</sup> for the globe, and were 4.1...9 (USA), 2.2... (China), 1.7... (EU27) and 0.5 (... [40])

### 3.1.4 Year 2024 Projection

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Globally, we estimate that global fossil CO<sub>2</sub> emissions (including cement carbonation, -0.21 GtC) will grow by 0.8% in 2024 (-0.3% to +1.9%) to 10.2 GtC (37.4 GtCO<sub>2</sub>), an historical record high<sup>2</sup>. Carbon Monitor projects a similar 2024 increase of 0.6% (-0.7% to 1.9%). GCB estimates of changes in 2024 emissions per fuel types, relative to 2023, are projected to be 0.2% (range -1.0% to 1.4%) for coal, +0.9% (range 0.0% to 1.8%) for oil, +2.4% (range 1.1% to 3.8%) for natural gas, and -2.8% (range -4.7% to -0.9%) for cement.

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**Deleted:** 1.1...8% in 2023 (0...024 (-0.3% to 2...1.9%) to 10.0... GtC (36.8...7.4 GtCO<sub>2</sub>), exceeding the pre-COVID19 2019 emission levels of 9.9 GtC (36.3 GtCO<sub>2</sub>). Global (... [41])

**Deleted:** 2023...024 emissions per fuel types, relative to 2023, are projected to be +1.1...2% (range -1.0... to 2...4%) for coal, +1.5...9% (range 0.6...% to 2.3...8%) for oil, +0.5...4% (range -0.9...1% to 1...8%) for natural gas, and 0...2.8% (range -0...7% to 2.4 (... [42])

For China, projected fossil emissions in 2024 are expected to increase slightly by 0.2% (range -1.6% to 2.0%) compared with 2023 emissions, bringing 2023 emissions for China around 3.3 GtC yr<sup>-1</sup> (12.0 GtCO<sub>2</sub> yr<sup>-1</sup>). In comparison, the Carbon Monitor estimate projects a 2024 decrease of 0.8% (range -3.8% to 1.9%). Our projected changes by fuel for China are +0.3% for coal, +0.8% for oil, +8.0% natural gas, and -8.1% for cement.

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For the USA, using the Energy Information Administration (EIA) emissions projection for 2024 combined with cement clinker data from USGS, we project a decrease of 0.6% (range -2.9% to 1.7%) compared to 2023, bringing USA 2023 emissions to around 1.3 GtC yr<sup>-1</sup> (4.9 GtCO<sub>2</sub> yr<sup>-1</sup>). Carbon Monitor projects a 2024 increase in USA emissions of 1.2% (-1.0% to 3.5%). Our projected changes by fuel are -3.5% for coal, -0.7% for oil, +1.0% for natural gas, and -5.8% for cement.

**Deleted:** 2023...024 combined with cement clinker data from USGS gives an... we project a decrease of 3...6% (range -5.0...9% to -...0...%) compared to 2022...023, bringing USA 2023 emissions to around 1.3 GtC yr<sup>-1</sup> (4.9 GtCO<sub>2</sub> yr<sup>-1</sup>). This is based on separate projections...arbon Monitor a 2024 increase in USA emissions of 1.2% (-1.0% to 3.5%). Our projected changes by fuel are -3.5% for coal -18.3%,... -0.7% for oil -... +1.0.3%,... for natural gas +1.4%,... and -5.8% for cement -4.0%. (... [44])

For the European Union, our projection for 2024 is for a decrease of 3.8% (range -6.2% to -1.4%) relative to 2023, with 2024 emissions around 0.7 GtC yr<sup>-1</sup> (2.4 GtCO<sub>2</sub> yr<sup>-1</sup>). The Carbon Monitor projection for the EU27 is slightly lower than GCB with a decrease of 5.5% (-9.2% to -1.9%). Our projected changes by fuel are -15.8% for coal, +0.2% for oil, -1.3% for natural gas, and -3.5% for cement.

**Deleted:** is for a decrease of 7.4% (range -9.9% to -4.9%) over 2022... with 2023...024 emissions around 0.7 GtC yr<sup>-1</sup> (2.6 GtCO<sub>2</sub> yr<sup>-1</sup>). This is based on separate projections... GtCO<sub>2</sub> yr<sup>-1</sup>). The Carbon Monitor projection for the EU27 is slightly lower than GCB with a decrease of 5.5% (-9.2% to -1.9%). Our projected changes by fuel are -15.8% for coal of -18.8%,... +0.2% for oil, -1.5%,... for natural gas -6.6%,... and -3.5% for cement -8.7%. (... [45])

For India, our projection for 2024 is an increase of 4.6% (range of 3.0% to 6.1%) over 2023, with 2024 emissions around 0.9 GtC yr<sup>-1</sup> (3.2 GtCO<sub>2</sub> yr<sup>-1</sup>). The Carbon Monitor projection for India is an increase of 5.5% (1.9% to 9.1%). Our projected changes by fuel are +4.5% for coal, +3.6% for oil, +11.8% for natural gas, and +4.0% for cement.

**Deleted:** 2023...024 is an increase of 8.2...6% (range of 6.7...0% to 9.7...1%) over 2022...023, with 2023...024 emissions around 0.8... GtC yr<sup>-1</sup> (3.1... GtCO<sub>2</sub> yr<sup>-1</sup>). This...he Carbon Monitor projection for India is based on separate projections...n increase of 5.5% (1.9% to 9.1%). Our projected changes by fuel are +4.5% for coal of +9.5%,... +3.6% for oil +5.3%,... +11.8% for natural gas +5.6%,... and +4.0% for cement +8.8%. (... [46])

International aviation and shipping are projected to increase by 7.8% in 2024, with international aviation projected to be up 14% over 2023, continuing to recover from pandemic lows, and international shipping projected to rise by 3%. The Carbon Monitor projects international aviation and shipping to increase by 3.3% in 2024.

**Deleted:** (2.8% of global emissions) ...re projected to increase by 11.9...8% in 2023...024, with ...nternational aviation projected to be up 28...4% over 2022...023, continuing to recover from pandemic lows, and international shipping projected to rise by 1%. (... [47])

<sup>2</sup> Growth rates in this section use a leap year adjustment that corrects for the extra day in 2024.

For the rest of the world, the expected change for 2024 is an increase of 1.1% (range -1.0% to 3.3%) with 2024 emissions around 4.0 GtC yr<sup>-1</sup> (14.5 GtCO<sub>2</sub> yr<sup>-1</sup>), similar to the Carbon Monitor projection of 1.1% (range -0.1% to 2.3%). The fuel-specific projected 2024 growth rates for the rest of the world are: +0.5% for coal, +0.5% for oil, +2.2% for natural gas, +2.0% for cement.

For traceability, Table S6 provides a comparison of annual projections from GCB since 2015 with the actual emissions assessed in the subsequent GCB annual report.

## 3.2 Emissions from Land Use Change

### 3.2.1 Historical period 1850-2023

Cumulative CO<sub>2</sub> emissions from land-use change (ELUC) for 1850-2023 were 225 ± 65 GtC (Table 8; Figure 3; Figure 16). The cumulative emissions from ELUC show a large spread among individual estimates of 150 GtC (H&C2023), 205 GtC (OSCAR), 250 GtC (LUCE) and 285 GtC (BLUE) for the four bookkeeping models and a similar wide estimate of 250 ± 85 GtC for the DGVMs (all cumulative numbers are rounded to the nearest 5 GtC). Vegetation biomass observations provide independent constraints on the ELUC estimates (Li et al., 2017). Over the 1901-2012 period, the GCB bookkeeping models cumulative ELUC amounts to 165 GtC [105 to 210 GtC], similar to the observation-based estimate of 155 ± 50 GtC (Li et al., 2017).

### 3.2.2 Recent period 1960-2023

In contrast to growing fossil emissions, CO<sub>2</sub> emissions from land-use, land-use change, and forestry remained relatively constant (around 1.5 GtC yr<sup>-1</sup>) over the 1960-1999 period. Since then, they have shown a statistically significant decrease of about 0.2 GtC per decade, reaching 1.3 ± 0.7 GtC yr<sup>-1</sup> for the 2014-2023 period (Table 7), but with significant spread, from 0.8 to 1.3 GtC yr<sup>-1</sup> across the four bookkeeping models (Table 5, Figure 7). Different from the bookkeeping average, the DGVMs average grows slightly larger over the 1980-2010 period and shows no sign of decreasing emissions in the recent decades, apart from in the most recent decade (Table 5, Figure 7). This is, however, expected as DGVM-based estimates include the loss of additional sink capacity, which grows with time, while the bookkeeping estimates do not (Supplement S.6.4).

We separate net ELUC into five component fluxes to gain further insight into the drivers of net emissions: deforestation, forest (re-)growth, wood harvest and other forest management, peat drainage and peat fires, and all other transitions (Figure 7c; supplemental Sec. S.2.1). We further decompose the deforestation and the forest (re-)growth term into contributions from shifting cultivation vs permanent forest cover changes (Figure 7d). Averaged over the 2014-2023 period and over the four bookkeeping estimates, fluxes from deforestation amount to 1.7 [1.4 to 2.3] GtC yr<sup>-1</sup> (Table 5), of which 1.0 [0.8, 1.1] GtC yr<sup>-1</sup> are from permanent deforestation. Fluxes from forest (re-)growth amount to -1.2 [-1.5, -0.9] GtC yr<sup>-1</sup> (Table 5), of which -0.5 [-0.7, -0.3] GtC yr<sup>-1</sup> are from re/afforestation and the remainder from forest regrowth in shifting cultivation cycles. Emissions from wood harvest and other forest management (0.3 [0.0, 0.6] GtC yr<sup>-1</sup>), peat drainage and peat fires (0.2 [0.2, 0.3] GtC yr<sup>-1</sup>) and the net flux from other transitions (0.1 [0.0, 0.1] GtC yr<sup>-1</sup>) are substantially less important globally (Table

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5). However, the small net flux from wood harvest and other forest management contains substantial gross fluxes that largely compensate each other (see Figure S8): 1.4 [0.9, 2.0] GtC yr<sup>-1</sup> emissions result from the decomposition of slash and the decay of wood products and -1.1 [-1.4, -0.8] GtC yr<sup>-1</sup> removals result from regrowth after wood harvesting.

The split into component fluxes clarifies the potentials for emission reduction and carbon dioxide removal: the emissions from permanent deforestation - the largest of our component fluxes - could be halted (largely) without compromising carbon uptake by forests, contributing substantially to emissions reduction. By contrast, reducing wood harvesting would have limited potential to reduce emissions as it would be associated with less forest regrowth; removals and emissions cannot be decoupled here on long timescales. A similar conclusion applies to removals and emissions from shifting cultivation, which we have therefore separated out. Carbon Dioxide Removal (CDR) in forests could instead be increased by permanently increasing the forest cover through re/afforestation. Our estimate of about -0.5 GtC yr<sup>-1</sup> removed on average each year during 2014-2023 by re/afforestation is similar to independent estimates that were derived from NGHGs for CDR in managed forests (through re/afforestation plus forest management) for 2013-2022 (-0.5 GtC yr<sup>-1</sup>, Pongratz et al., 2024). Re/afforestation constitutes the vast majority of all current CDR (Pongratz et al., 2024). Though they cannot be compared directly to annual fluxes from the atmosphere, CDR through transfers between non-atmospheric reservoirs such as in durable HWP, biochar, or BECCS comprise much smaller amounts of carbon. 218 MtC yr<sup>-1</sup> have been estimated to be transferred to HWP, averaged over 2013-2022. The net flux of HWP, considering the re-release of CO<sub>2</sub> through their decay, amounts to 91 MtC yr<sup>-1</sup> over that period (Pongratz et al., 2024). Note that some double-counting between the CDR through HWP and the CDR through re/afforestation exists if the HWP are derived from newly forested areas. BECCS projects have been estimated to store 0.1 MtC yr<sup>-1</sup> in geological projects worldwide in 2023, biochar projects 0.2 MtC yr<sup>-1</sup> (Pongratz et al., 2024). “Blue carbon”, i.e. coastal wetland management such as restoration of mangrove forests, saltmarshes and seagrass meadows, though at the interface of land and ocean carbon fluxes, are counted towards the land-use sector as well. Currently, bookkeeping models do not include blue carbon; however, current CDR deployment in coastal wetlands is small globally, less than 0.003 MtC yr<sup>-1</sup> (Powis et al., 2023).

The statistically significant decrease in E<sub>LUC</sub> since the late-1990s, including the larger drop within the most recent decade, is due to the combination of decreasing emissions from deforestation (in particular permanent deforestation) and increasing removals from forest regrowth (with those from re/afforestation stagnating globally in the last decade). Emissions in 2014-2023 are 28% lower than in the late-1990s (1995-2004) and 20% lower than in 2004-2013. The steep drop in E<sub>LUC</sub> after 2015 is due to the combined effect from a peak in peat fire emissions in 2015 and a long-term decline in deforestation emissions in many countries over the 2010-2020 period with largest declines in the Democratic Republic of the Congo, Brazil, China, and Indonesia. Since the processes behind gross removals, foremost forest regrowth and soil recovery, are all slow, while gross emissions include a large instantaneous component, short-term changes in land-use dynamics, such as a temporary decrease in deforestation, influences gross emissions dynamics more than gross removals dynamics, which rather are a response to longer-term dynamics. Component fluxes often differ more across the four bookkeeping estimates than the net flux, which is expected due to different process representation; in particular, the treatment

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of shifting cultivation, which increases both gross emissions and removals, differs across models, but also net and gross wood harvest fluxes show high uncertainty. By contrast, models agree relatively well for emissions from permanent deforestation.

Overall, highest land-use emissions occur in the tropical regions of all three continents. The top three emitters (both cumulatively 1959-2023 and on average over 2014-2023) are Brazil (in particular the Amazon Arc of Deforestation), Indonesia and the Democratic Republic of the Congo, with these 3 countries contributing 0.7 GtC yr<sup>-1</sup> or 60% of the global net land-use emissions (average over 2014-2023) (Figure 6b, Figure 7b). This is related to massive expansion of cropland, particularly in the last few decades in Latin America, Southeast Asia, and sub-Saharan Africa (Hong et al., 2021), to a substantial part for export of agricultural products (Pendrill et al., 2019). Emission intensity is high in many tropical countries, particularly of Southeast Asia, due to high rates of land conversion in regions of carbon-dense and often still pristine, undegraded natural forests (Hong et al., 2021). Emissions are further increased by peat fires in equatorial Asia (GFED4s, van der Werf et al., 2017). Our estimates of high ELUC in China has been revised down since the 1980s as compared to GCB2023 related to the update of the land-use forcing, which is now based on the cropland dataset by Yu et al. (2022) (see Supplement S.2.2), which suggests lower cropland expansion and thus less deforestation than the previous datasets assumed. Uptake due to land-use change occurs in several regions of the world (Figure 6b) particularly because of re/afforestation. Highest CDR in the last decade is seen in China, where our estimates show an even larger uptake since 2010 compared to GCB2023 related to the updated land-use forcing, in the EU27, partly related to expanding forest area as a consequence of the forest transition in the 19<sup>th</sup> and 20<sup>th</sup> century and subsequent regrowth of forest (Mather 2001; McGrath et al., 2015), and in the U.S. Substantial uptake through re/afforestation also exists in other regions such as Brazil, Myanmar or Russia, where, however, emissions from deforestation and other land-use changes dominate the net flux.

While the mentioned patterns are robust and supported by independent literature, we acknowledge that model spread is substantially larger on regional than global levels, as has been shown for bookkeeping models (Bastos et al., 2021) as well as DGVMs (Obermeier et al., 2021). Assessments for individual regions are being performed as part of REgional Carbon Cycle Assessment and Processes (RECCAP2; Ciais et al., 2020, Poulter et al., 2022) or already exist for selected regions (e.g., for Europe by Petrescu et al., 2020, for Brazil by Rosan et al., 2021, for 8 selected countries/regions in comparison to inventory data by Schwingshackl et al., 2022). The revisions since GCB2023 reflect such uncertainties: The integration of a fourth bookkeeping model alters our estimates, though only to a limited extent given that the new model LUCE lies in between the other three models for the global ELUC estimates. Larger changes are obvious at regional level due to the revisions of the land-use forcing with a general update to more recent FAO input for agricultural areas and wood harvest, new MapBiomass input for Brazil and Indonesia and the updated cropland dataset in China.

The NGHGI data under the LULUCF sector and the LULUCF estimates from FAOSTAT differ from the global models' definition of ELUC (see Section 2.2.1). In the NGHGI reporting, the natural fluxes (S<sub>LAND</sub>) are counted towards ELUC when they occur on managed land (Grassi et al., 2018). To compare our results to the NGHGI approach, we perform a translation of our ELUC estimates by adding S<sub>LAND</sub> in managed forest from the DGVMs

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simulations (following the methodology described in Grassi et al., 2023) to the bookkeeping  $E_{LUC}$  estimate (see Supplement S.2.3). For the 2014-2023 period, we estimate that  $1.8 \text{ GtC yr}^{-1}$  of  $S_{LAND}$  occurred in managed forests. Adding this sink to  $E_{LUC}$  changes  $E_{LUC}$  from being a source of  $1.1 \text{ GtC yr}^{-1}$  to a sink of  $0.7 \text{ GtC yr}^{-1}$ , very similar to the NGHGI estimate that yields a sink of  $0.8 \text{ GtC yr}^{-1}$  (Figure 8, Table S10). We further apply a mask of managed land to the net atmosphere-to-land flux estimate from atmospheric inversions to obtain inverse estimates that are comparable to the NGHGI estimates and to the translated  $E_{LUC}$  estimates from bookkeeping models (see Supplement S.2.3). The inversion-based net flux in managed land indicates a sink of  $0.7 \text{ GtC yr}^{-1}$  for 2014-2023, which agrees very well with the NGHGI and the translated  $E_{LUC}$  estimates (Figure 8, Table S10). Additionally, the interannual variability of the inversion estimates and the translated  $E_{LUC}$  estimates show a remarkable agreement (Pearson correlation of 0.81 in 2000-2023), which supports the suggested translation approach.

The translation approach has been shown to be generally applicable also at the country-level (Grassi et al., 2023; Schwingshackl et al., 2022). Country-level analysis suggests, e.g., that the bookkeeping method estimates higher deforestation emissions than the national report in Indonesia, but less  $\text{CO}_2$  removal by afforestation than the national report in China. The fraction of the natural  $\text{CO}_2$  sinks that the NGHGI estimates include differs substantially across countries, related to varying proportions of managed vs total forest areas (Schwingshackl et al., 2022). By comparing  $E_{LUC}$  and NGHGI on the basis of the component fluxes used above, we find that our estimates reproduce very closely the NGHGI estimates for emissions from permanent deforestation, peat emissions, and other transitions (Figure 8), although a difference in sign (small source in bookkeeping estimates, small sink in NGHGI) creates a notable difference between NGHGI and bookkeeping estimates. Fluxes due to forest (re-)growth & other forest management, that is, (re-)growth from re/afforestation plus the net flux from wood harvesting and other forest management and emissions and removals in shifting cultivation cycles, constitute a large sink in the NGHGI ( $-1.9 \text{ GtC yr}^{-1}$  averaged over 2014-2023), since they also include  $S_{LAND}$  in managed forests. Summing up the bookkeeping estimates of (re-)growth from re/afforestation, the net flux from wood harvesting and other forest management, and the emissions and removals in shifting cultivation cycles, and adding  $S_{LAND}$  in managed forests yields a flux of  $-2.0 \text{ GtC yr}^{-1}$  (averaged over 2014-2023), which compares well with the NGHGI estimate. Though estimates between NGHGI, FAOSTAT and the translated budget estimates still differ in value and need further analysis, the approach suggested by Grassi et al. (2023), which we adopt here, provides a feasible way to relate the global models' and NGHGI approach to each other and thus link the anthropogenic carbon budget estimates of land  $\text{CO}_2$  fluxes directly to the Global Stocktake, as part of the UNFCCC Paris Agreement.

### 3.2.3 Final year 2023

The global  $\text{CO}_2$  emissions from land-use change are estimated as  $1.0 \pm 0.7 \text{ GtC}$  in 2023, similar to the 2022 estimate. However, confidence in the annual change remains low. Despite El Niño conditions, which in general lead to more fires in deforestation areas, peat fire emissions in Indonesia remained below average (GFED4.1s; updated from van der Werf et al., 2017). In South America, emissions from tropical deforestation and

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2592 degradation fires have been about average, as effects of the El Niño in the Amazon, such as droughts, are not  
2593 expected before 2024.

### 2594 3.2.4 Year 2024 Projection

2595 In Southeast Asia, peat fire emissions have further dropped (from 27 Tg C in 2023 to 1 Tg C in 2024 through  
2596 October 17 2024; GFED4.1s, van der Werf et al., 2017), as have tropical deforestation and degradation fires  
2597 (from 33 Tg C to 6 Tg C) as the El Niño conditions ceased. By contrast, emissions from tropical deforestation  
2598 and degradation fires in South America have risen from 121 Tg C in 2023 to 324 Tg C in 2024 up until October  
2599 17, as the impacts of the El Niño unfold, in particular drought conditions since 2023. The 2024 South American  
2600 fire emissions are among the highest values in the record, which started in 1997. Part of the increase is due to  
2601 elevated fire activity in the wetlands of the Pantanal. Disentangling the degree to which interannual variability in  
2602 rainfall patterns and stronger environmental protection measures in both Indonesia after their 2015 high fire  
2603 season and in Brazil after the change in government play a role in fire trends is an important research topic.  
2604 Cumulative 2024 fire emission estimates through October 17 2024 are 422 Tg C for global deforestation and  
2605 degradation fires and 1 Tg C for peatland fires in Southeast Asia.

2606 Based on these estimates, we expect  $E_{LUC}$  emissions of around 1.1 GtC ( $4.2 \text{ GtCO}_2$ ) in 2024, slightly above the  
2607 2023 level. Note that although our extrapolation includes tropical deforestation and degradation fires, the  
2608 degradation attributable to selective logging, edge-effects or fragmentation is not captured. Further,  
2609 deforestation and fires in deforestation zones may become more disconnected, partly due to changes in  
2610 legislation in some regions. For example, Van Wees et al. (2021) found that the contribution from fires to forest  
2611 loss decreased in the Amazon and in Indonesia over the period of 2003-2018.

### 2612 3.3 CDR not based on vegetation

2613 Besides the CDR through land use (Sec. 3.2), the atmosphere to geosphere flux of carbon resulting from carbon  
2614 dioxide removal (CDR) activity in 2023 is estimated at 0.011 MtC/yr. This results primarily from 0.009 MtC/yr  
2615 of enhanced weathering projects and 0.001 MtC/yr of DACCS. While it represents a growth of 200% in the  
2616 anthropogenic sink, from the 0.0036 MtC/yr estimate in 2022, it remains about a million times smaller than  
2617 current fossil  $\text{CO}_2$  emissions.

### 2618 3.4 Total anthropogenic emissions

2619 Cumulative anthropogenic  $\text{CO}_2$  emissions (fossil and land use) for 1850-2023 totalled 710  $\pm 70$  GtC (2605  $\pm$   
2620  $260 \text{ GtCO}_2$ ), of which 70% (500 GtC) occurred since 1960 and 34% (240 GtC) since 2000 (Table 7 and 8).  
2621 Total anthropogenic emissions more than doubled over the last 60 years, from  $4.6 \pm 0.7 \text{ GtC yr}^{-1}$  for the decade  
2622 of the 1960s to an average of  $10.8 \pm 0.9 \text{ GtC yr}^{-1}$  during 2014-2023, and reaching  $11.1 \pm 0.9 \text{ GtC}$  (40.5  $\pm 3.2$   
2623  $\text{GtCO}_2$ ) in 2023. However, total anthropogenic  $\text{CO}_2$  emissions have been stable over the last decade (zero  
2624 growth rate over the 2014-2023 period), much slower than the 2.0% growth rate over the previous decade  
2625 (2004-2013).

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Deleted: 1...  $\text{GtCO}_2$ ) in 2023. Our preliminary estimate of  $E_{LUC}$  for 2023 is substantially lower than...024, slightly the 2013-2022 average, which saw years of anomalously dry conditions in Indonesia and high deforestation fires in South America (Friedlingstein et al., 2022b). (... [56])

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During the historical period 1850-2023, 31% of historical emissions were from land use change and 69% from fossil emissions. However, fossil emissions have grown significantly since 1960 while land use changes have not, and consequently the contributions of land use change to total anthropogenic emissions were smaller during recent periods, 18% during the period 1960-2023 and down to 10% over the last decade (2014-2023). For 2024, we project global total anthropogenic CO<sub>2</sub> emissions from fossil and land use changes to be around 11.3 GtC (41.6 GtCO<sub>2</sub>), 2% above the 2023 level. All values here include the cement carbonation sink (currently about 0.2 GtC yr<sup>-1</sup>).

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### 3.5 Atmospheric CO<sub>2</sub>

#### 3.5.1 Historical period 1850-2023

Atmospheric CO<sub>2</sub> concentration was approximately 278 parts per million (ppm) in 1750, reaching 300 ppm in the late 1900s, 350 ppm in the late 1980s, and reaching 419.31 ± 0.1 ppm in 2023 (Lan et al., 2024; Figure 1). The mass of carbon in the atmosphere increased by 51% from 590 GtC in 1750 to 890 GtC in 2023. Current CO<sub>2</sub> concentrations in the atmosphere are unprecedented in the last 2 million years and the current rate of atmospheric CO<sub>2</sub> increase is at least 10 times faster than at any other time during the last 800,000 years (Canadell et al., 2021).

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#### 3.5.2 Recent period 1960-2023

The growth rate in atmospheric CO<sub>2</sub> level increased from 1.7 ± 0.07 GtC yr<sup>-1</sup> in the 1960s to 5.2 ± 0.02 GtC yr<sup>-1</sup> during 2014-2023 with important decadal variations (Table 7, Figure 3 and Figure 4). During the last decade (2014-2023), the growth rate in atmospheric CO<sub>2</sub> concentration continued to increase, albeit with large interannual variability (Figure 4).

The airborne fraction (AF) is defined as the ratio of atmospheric CO<sub>2</sub> growth rate to total anthropogenic emissions:

$$AF = G_{ATM} / (E_{FOS} + E_{LUC}) \quad (2)$$

It provides a diagnostic of the relative strength of the land and ocean carbon sinks in removing part of the anthropogenic CO<sub>2</sub> perturbation. The evolution of AF over the last 60 years shows no significant trend, remaining at around 44%, albeit showing a large interannual and decadal variability driven by the year-to-year variability in G<sub>ATM</sub> (Figure 10). The observed stability of the airborne fraction over the 1960-2023 period indicates that the ocean and land CO<sub>2</sub> sinks have been increasing in pace with the total anthropogenic emissions over that period, removing on average about 56% of the emissions (see Sections 3.6.2 and 3.7.2).

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2752	<b>3.5.3 Final year 2023</b>	Deleted: 2022
2753	The growth rate in atmospheric CO <sub>2</sub> concentration was $5.9 \pm 0.2$ GtC ( $2.79 \pm 0.08$ ppm) in 2023 (Figure 4; Lan	Formatted: Outline numbered + Level: 3 + Numbering Style: 1, 2, 3, ... + Start at: 1 + Alignment: Left + Aligned at: 0 cm + Indent at: 1.27 cm
2754	et al., 2024), well above the 2022 growth rate ( $4.6 \pm 0.2$ GtC) or the 2014-2023 average ( $5.2 \pm 0.02$ GtC), as to	Deleted: 4.6...9 $\pm 0.2$ GtC ( $2.18...9 \pm 0.08$ ppm) in 2022...023 (Figure 4; Lan et al., 2023), below...024), well
2755	be expected during an El Niño year. The 2023 atmospheric CO <sub>2</sub> growth rate was the 4th largest over the 1959-	above the 2021...022 growth rate ( $5.2...6 \pm 0.2$ GtC) or the
2756	2023 atmospheric observational record, closely following 2015, 2016 and 1998, all strong El Niño years.	2013-2022...014-2023 average ( $5.2 \pm 0.02$ GtC). ... [59]
2757	<b>3.5.4 Year 2024 Projection</b>	Deleted: 2023
2758	The 2024 growth in atmospheric CO <sub>2</sub> concentration (G <sub>ATM</sub> ) is projected to be about $5.9$ GtC ( $2.76$ ppm), still	Formatted: Outline numbered + Level: 3 + Numbering Style: 1, 2, 3, ... + Start at: 1 + Alignment: Left + Aligned at: 0 cm + Indent at: 1.27 cm
2759	high, which is common for the year after a strong El Niño year. This is the average of the GCB regression	Deleted: 2023...024 growth in atmospheric CO <sub>2</sub>
2760	method ( $5.6$ GtC, $2.64$ ppm) and ESMs the multi-model mean ( $6.1$ GtC, $2.88$ ppm). The 2024 atmospheric CO <sub>2</sub>	concentration (G <sub>ATM</sub> ) is projected to be about $5.1... \pm 0.1$ GtC
2761	concentration, averaged over the year, is expected to reach the level of $422.5$ ppm, $52\%$ over the pre-industrial	( $2.4...6$ ppm)..., still high, which is common for the year
2762	level.	after a strong El Niño year. This is the average of the GCB
2763	<b>3.6 Ocean Sink</b>	regression method ( $5.07... \pm 0.1$ GtC, $2.39...4$ ppm) and ESMs the
2764	<b>3.6.1 Historical period 1850-2023</b>	multi-model mean ( $5.11...1 \pm 0.1$ GtC, $2.41...8$ ppm). The
2765	Cumulated since 1850, the ocean sink adds up to $185 \pm 35$ GtC, with more than two thirds of this amount ( $130 \pm$	2023...024 atmospheric CO <sub>2</sub> concentration, averaged over
2766	$25$ GtC) being taken up by the global ocean since 1960. Over the historical period, the ocean sink increased in	the year, is expected to reach the level of $419.3...22.5$ ppm,
2767	pace with the anthropogenic emissions exponential increase (Figure 3). Since 1850, the ocean has removed 26%	51 ... [60]
2768	of total anthropogenic emissions.	Formatted: Indent: Left: 0 cm, First line: 0 cm, Outline numbered + Level: 2 + Numbering Style: 1, 2, 3, ... + Start at: 1 + Alignment: Left + Aligned at: 0 cm + Indent at: 1.02 cm
2769	<b>3.6.2 Recent period 1960-2023</b>	Deleted: 2022
2770	The ocean CO <sub>2</sub> sink increased from $1.2 \pm 0.4$ GtC yr <sup>-1</sup> in the 1960s to $2.9 \pm 0.4$ GtC yr <sup>-1</sup> during 2014-2023	Formatted: Outline numbered + Level: 3 + Numbering Style: 1, 2, 3, ... + Start at: 1 + Alignment: Left + Aligned at: 0 cm + Indent at: 1.27 cm
2771	(Table 7), with interannual variations of the order of a few tenths of GtC yr <sup>-1</sup> (Figure 4, Figure 11). The ocean-	Deleted: 180...85 $\pm 35$ GtC, with more than two thirds of this
2772	borne fraction (SO <sub>OCEAN</sub> /(E <sub>FOS</sub> +E <sub>LUC</sub> )) has been remarkably constant around 25% on average (Figure 10c), with	amount (125 ... [61]
2773	variations around this mean illustrating the decadal variability of the ocean carbon sink. So far, there is no	Deleted: 2022
2774	indication of a decrease in the ocean-borne fraction from 1960 to 2022. The increase of the ocean sink is	Formatted: Outline numbered + Level: 3 + Numbering Style: 1, 2, 3, ... + Start at: 1 + Alignment: Left + Aligned at: 0 cm + Indent at: 1.27 cm
2775	primarily driven by the increased atmospheric CO <sub>2</sub> concentration, with the strongest CO <sub>2</sub> induced signal in the	Deleted: 1... $\pm 0.4$ GtC yr <sup>-1</sup> in the 1960s to $2.8... \pm 0.4$ GtC
2776	North Atlantic and the Southern Ocean (Figure 12a). The effect of climate change is much weaker, reducing the	yr <sup>-1</sup> during 2013-2022...014-2023 (Table 7), with interannual
2777	ocean sink globally by $0.17 \pm 0.05$ GtC yr <sup>-1</sup> ( $-5.9\%$ of SO <sub>OCEAN</sub> ) during 2014-2023 (all models simulate a	variations of the order of a few tenths of GtC yr <sup>-1</sup> (Figure
2778	weakening of the ocean sink by climate change, range $-3.4$ to $-10.7\%$ ), and does not show clear spatial patterns	10..., Figure 11). The ocean-borne fraction
2779	across the GOBMs ensemble (Figure 12b). This is the combined effect of change and variability in all	(SO <sub>OCEAN</sub> /(E <sub>FOS</sub> +E <sub>LUC</sub> )) has been remarkably constant around
2780	atmospheric forcing fields, previously attributed to wind and temperature changes (LeQuéré et al., 2010, Bunsen	25% on average (Figure 9c...0c), with variations around this
2781	et al., 2024). The effect of warming is smaller than expected from offline calculation due to a stabilising	mean illustrating the decadal variability of the ocean carbon
2782	feedback from limited exchange between surface and deep waters (Bunsen et al., 2024).	sink. So far, there is no indication of a decrease in the ocean-
2783	The global net air-sea CO <sub>2</sub> flux is a residual of large natural and anthropogenic CO <sub>2</sub> fluxes into and out of the	borne fraction from 1960 to 2022. The increase of the ocean
2784	ocean with distinct regional and seasonal variations (Figure 6 and S1). Natural fluxes dominate on regional	sink is primarily driven by the increased atmospheric CO <sub>2</sub>
		concentration, with the strongest CO <sub>2</sub> induced signal in the
		North Atlantic and the Southern Ocean (Figure 11a...2a). The
		effect of climate change is much weaker, reducing the ocean
		sink globally by $0.16...7 \pm 0.04...5$ GtC yr <sup>-1</sup> ( $-6.7...9\%$ of
		SO <sub>OCEAN</sub> ) during 2013-2022...014-2023 (all models simulate a
		weakening of the ocean sink by climate change, range $-$
		$3.4.3...to -10.3...%$ ), and does not show clear spatial patterns
		across the GOBMs ensemble (Figure 11b...2b). This is the
		combined effect of change and variability in all atmospheric
		forcing fields, previously attributed, in one model, ... [62]
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scales, but largely cancel out when integrated globally (Gruber et al., 2009). Mid-latitudes in all basins and the high-latitude North Atlantic dominate the ocean CO<sub>2</sub> uptake where low temperatures and high wind speeds facilitate CO<sub>2</sub> uptake at the surface (Takahashi et al., 2009). In these regions, formation of mode, intermediate and deep-water masses transport anthropogenic carbon into the ocean interior, thus allowing for continued CO<sub>2</sub> uptake at the surface. Outgassing of natural CO<sub>2</sub> occurs mostly in the tropics, especially in the equatorial upwelling region, and to a lesser extent in the North Pacific and polar Southern Ocean, mirroring a well-established understanding of regional patterns of air-sea CO<sub>2</sub> exchange (e.g., Takahashi et al., 2009, Gruber et al., 2009). These patterns are also noticeable in the Surface Ocean CO<sub>2</sub> Atlas (SOCAT) dataset, where an ocean fCO<sub>2</sub> value above the atmospheric level indicates outgassing (Figure S1). This map further illustrates the data-sparsity in the Indian Ocean and the southern hemisphere in general.

The largest variability in the ocean sink occurs on decadal time-scales (Figure 11). The ensemble means of GOBMs and fCO<sub>2</sub>-products show the same patterns of decadal variability, although with a larger amplitude of variability in the fCO<sub>2</sub>-products than in the GOBMs. The ocean sink stagnated in the 1990s and strengthened between the early 2000s and the mid-2010s (Figure 11; Le Quéré et al., 2007; Landschützer et al., 2015, 2016; DeVries et al., 2017; Hauck et al., 2020; McKinley et al., 2020, Gruber et al., 2023). More recently, the sink seems to have entered a phase of stagnation since 2016, largely in response to large inter-annual climate variability. Different explanations have been proposed for the decadal variability in the 1990s and 2000s, ranging from the ocean's response to changes in atmospheric wind systems (e.g., Le Quéré et al., 2007, Keppler and Landschützer, 2019), including variations in upper ocean overturning circulation (DeVries et al., 2017) to the eruption of Mount Pinatubo in the 1990s (McKinley et al., 2020). The main origin of the decadal variability is a matter of debate with a number of studies initially pointing to the Southern Ocean (see review in Canadell et al., 2021), but also contributions from the North Atlantic and North Pacific (Landschützer et al., 2016, DeVries et al., 2019), or a global signal (McKinley et al., 2020) were proposed.

On top of the decadal variability, interannual variability of the ocean carbon sink is driven by climate variability with a first-order effect from a stronger ocean sink during large El Niño events (e.g., 1997-1998) (Figure 11; Rödenbeck et al., 2014, Hauck et al., 2020; McKinley et al. 2017) leading to a reduction in CO<sub>2</sub> outgassing from the Tropical Pacific. During 2010-2016, the ocean CO<sub>2</sub> sink appears to have intensified in line with the expected increase from atmospheric CO<sub>2</sub> (McKinley et al., 2020). This effect is similar in the fCO<sub>2</sub>-products (Figure 11, ocean sink 2016 minus 2010, GOBMs:  $+0.42 \pm 0.11$  GtC yr<sup>-1</sup>, fCO<sub>2</sub>-products:  $+0.44$  GtC yr<sup>-1</sup>, range 0.18 to 0.72 GtC yr<sup>-1</sup>). The reduction of  $-0.18$  GtC yr<sup>-1</sup> (range:  $-0.41$  to  $-0.03$  GtC yr<sup>-1</sup>) in the ocean CO<sub>2</sub> sink in 2017 is consistent with the return to normal conditions after the El Niño in 2015/16, which caused an enhanced sink in previous years. After an increasing SO<sub>CEAN</sub> in 2018 and 2019, the GOBM and fCO<sub>2</sub>-product ensemble means suggest a decrease of SO<sub>CEAN</sub>, related to the triple La Niña event 2020-2022, followed by a rebound in 2023 linked to the onset of an El Niño event.

Although all individual GOBMs and fCO<sub>2</sub>-products fall within the observational constraint, the ensemble means of GOBMs, and fCO<sub>2</sub>-products (adjusted for the riverine flux) show a mean offset increasing from 0.31 GtC yr<sup>-1</sup> in the 1990s to 0.49 GtC yr<sup>-1</sup> in the decade 2014-2023 and a slightly lower offset of 0.3 GtC yr<sup>-1</sup> in 2023. In this

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version of the GCB, the  $S_{OCEAN}$  positive trend diverges over time by a factor of 1.4 since 2002 (GOBMs:  $0.25 \pm 0.04$  GtC yr<sup>-1</sup> per decade,  $fCO_2$ -products:  $0.35$  GtC yr<sup>-1</sup> per decade,  $[0.17$  to  $0.79$  GtC yr<sup>-1</sup> per decade],  $S_{OCEAN}$ :  $0.30$  GtC yr<sup>-1</sup> per decade), but the uncertainty ranges overlap. This divergence is smaller than reported in previous GCB versions, because of the updated lower sink estimates by the  $fCO_2$ -products for recent years. This also leads to agreement on the trend since 2010 (GOBMs:  $0.18 \pm 0.06$  GtC yr<sup>-1</sup> per decade,  $fCO_2$ -products:  $0.18$  GtC yr<sup>-1</sup> per decade,  $[-0.36$  to  $0.73$  GtC yr<sup>-1</sup> per decade]  $S_{OCEAN}$ :  $0.18$  GtC yr<sup>-1</sup> per decade). A hybrid approach recently constrained the trend 2000-2022 to  $0.42 \pm 0.06$  GtC yr<sup>-1</sup> decade<sup>-1</sup> (Mayot et al., 2024), which aligns with the updated trends of  $S_{OCEAN}$  ( $0.39$  GtCyr<sup>-1</sup> decade<sup>-1</sup>) and of the  $fCO_2$ -products ( $0.45$   $[0.28, 0.84]$  GtCyr<sup>-1</sup> decade<sup>-1</sup>), while the GOBMs result in a lower trend ( $0.32 \pm 0.04$  GtC yr<sup>-1</sup> per decade) over the same period.

In the current dataset, the discrepancy between the two types of estimates stems from a persistently larger  $S_{OCEAN}$  in the  $fCO_2$ -products in the northern extra-tropics since around 2002 and an intermittently larger  $S_{OCEAN}$  in the southern extra-tropics in the period 2008-2020 (Figure J4). Note that the discrepancy in the mean flux, which was located in the Southern Ocean in GCB 2022 and earlier, was reduced due to the choice of the regional river flux adjustment (Lacroix et al., 2020 instead of Aumont et al., 2001). This comes at the expense of a discrepancy in the mean  $S_{OCEAN}$  of about  $0.2$  GtC yr<sup>-1</sup> in the tropics. Likely explanations for the discrepancy in the trends and decadal variability in the high-latitudes are data sparsity and uneven data distribution (Bushinsky et al., 2019, Gloege et al., 2021, Hauck et al., 2023a, Mayot et al., 2024). In particular, two  $fCO_2$ -products were shown to overestimate the Southern Ocean  $CO_2$  flux trend by 50 and 130% based on current sampling in a model subsampling experiment (Hauck et al., 2023a) and the largest trends in the  $fCO_2$ -products occurred in a data void region in the North Pacific (Mayot et al., 2024). In this respect it is highly worrisome that the coverage of  $fCO_2$  observations continues to decline (Dong et al 2024), and is now down to that of the early 2000s (Fig. 11). Another likely contributor to the discrepancy between GOBMs and  $fCO_2$ -products are model biases (as indicated by the comparison with Mayot et al., 2024, by the large model spread in the South, Figure J4, and the larger model-data  $fCO_2$  mismatch, Figure S2).

The reported  $S_{OCEAN}$  estimate from GOBMs and  $fCO_2$ -products is  $2.2 \pm 0.4$  GtC yr<sup>-1</sup> over the period 1994 to 2007, which is in agreement with the ocean interior estimate of  $2.2 \pm 0.4$  GtC yr<sup>-1</sup>, which accounts for the climate effect on the natural  $CO_2$  flux of  $-0.4 \pm 0.24$  GtC yr<sup>-1</sup> (Gruber et al., 2019) to match the definition of  $S_{OCEAN}$  used here (Hauck et al., 2020). This comparison depends critically on the estimate of the climate effect on the natural  $CO_2$  flux, which is smaller from the GOBMs ( $-0.1$  GtC yr<sup>-1</sup>) than in Gruber et al. (2019). Uncertainties of these two estimates would also overlap when using the GOBM estimate of the climate effect on the natural  $CO_2$  flux. Similarly, the  $S_{OCEAN}$  estimates integrated over the decades 1994-2004 ( $21.5$  GtC yr<sup>-1</sup>) and 2004-2014 ( $25.6$  GtC yr<sup>-1</sup>) agree with the interior ocean-based estimates of Müller et al. (2023;  $21.4 \pm 2.8$  and  $26.5 \pm 1.3$  GtC yr<sup>-1</sup>), but depend critically on assumptions of the climate effect on natural carbon, which in turn, are based on the  $fCO_2$ -products in Müller et al. (2023).

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In previous GCB releases, the ocean sink 1959-1989 was only estimated by GOBMs due to the absence of  $fCO_2$  observations. Now, the first data-based estimates extending back to 1957/58 are becoming available (Jena-MLS, Rödenbeck et al., 2022, LDEO-HPD, Bennington et al., 2022; Gloege et al., 2022). These are based on a multi-linear regression of  $pCO_2$  with environmental predictors (Rödenbeck et al., 2022) or on model-data  $pCO_2$  misfits and their relation to environmental predictors (Bennington et al., 2022). The Jena-MLS and LDEO-HPD estimates fall well within the range of GOBM estimates and have a correlation of 0.99 and 0.98 respectively with  $S_{OCEAN}$  for the period 1959-2022 (and 0.98 and 0.97 for the 1959-1989 period). They agree well on the mean  $S_{OCEAN}$  estimate since 1977 with a slightly higher amplitude of variability (Figure 10). Until 1976, Jena-MLS and LDEO-HPD are respectively abo... [63]

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### 3.6.3 Final year 2023

The estimated ocean CO<sub>2</sub> sink is  $2.9 \pm 0.4$  GtC for 2023. This is a small increase of  $0.16$  GtC compared to 2022, in line with the expected sink strengthening from the 2023 El Niño conditions. GOBM and fCO<sub>2</sub>-product ensemble mean estimates consistently result in an SOCEAN increase in 2023 (GOBMs:  $0.17 \pm 0.15$  GtC, fCO<sub>2</sub>-products:  $0.14 [-0.04, 0.30]$  GtC). Eight GOBMs and six fCO<sub>2</sub>-products show an increase in SOCEAN, while only two GOBMs and two fCO<sub>2</sub>-products show a minor decrease in SOCEAN of less than  $0.05$  GtC (Figure 11). The fCO<sub>2</sub>-products have a larger uncertainty at the end of the reconstructed time series, potentially linked to uncertainties related to fewer available observations in the final year and the shift from La Niña to El Niño (see e.g. Watson et al. 2020, Pérez et al. 2024). Specifically, the fCO<sub>2</sub>-products' estimate of the last year is regularly adjusted in the following release owing to the tail effect and an incrementally increasing data availability. While the monthly grid cells covered may have a lag of only about a year (Figure 11 inset), the values within grid cells may change with 1-5 years lag (see absolute number of observations plotted in previous GCB releases), potentially resulting in annual changes in the flux magnitude from fCO<sub>2</sub>-products.

### 3.6.4 Year 2024 Projection

Using a feed-forward neural network method (see Section 2.5.2) we project an ocean sink of  $3.0$  GtC for 2024, only  $0.1$  GtC higher than for the year 2023, consistent with El Niño to neutral conditions in 2024. The set of ESMs predictions support this estimate with a 2024 ocean sink of around  $3.0 [2.9, 3.1]$  GtC.

### 3.6.5 Evaluation of Ocean Models and fCO<sub>2</sub>-products

The process-based model evaluation draws a generally positive picture with GOBMs scattered around the observational values for Southern Ocean sea-surface salinity, Southern Ocean stratification index and surface ocean Revelle factor (Section 3.3 and Table S11). However, the Atlantic Meridional Overturning Circulation at 26°N is underestimated by 8 out of 10 GOBMs, and overestimated by one GOBM. It is planned to derive skill scores for the GOBMs in future releases based on these metrics.

The model simulations allow to separate the anthropogenic carbon component (steady state and non-steady state, sim D - sim A) and to compare the model flux and DIC inventory change directly to the interior ocean estimate of Gruber et al. (2019) without further assumptions (Table S11). The GOBMs ensemble average of anthropogenic carbon inventory changes 1994-2007 amounts to  $2.4$  GtC yr<sup>-1</sup> and is thus lower than the  $2.6 \pm 0.3$  GtC yr<sup>-1</sup> estimated by Gruber et al. (2019) although within the uncertainty. Only three models fall within the range reported by Gruber et al. (2019). This suggests that the majority of the GOBMs underestimate anthropogenic carbon uptake by 10-20% and some models even more. Comparison to the decadal estimates of anthropogenic carbon accumulation (Müller et al., 2023) are close to the interior ocean data based estimate for the decade 2004-2014 (GOBMs sim D minus sim A,  $24.7 \pm 3.6$  GtC yr<sup>-1</sup>, Müller et al.  $27.3 \pm 2.5$  GtC yr<sup>-1</sup>), but do not reproduce the supposedly higher anthropogenic carbon accumulation in the earlier period 1994-2004 (GOBMs sim D minus sim A,  $21.1 \pm 3.0$  GtC yr<sup>-1</sup>, Müller et al.  $29.3 \pm 2.5$  GtC yr<sup>-1</sup>). Analysis of Earth System Models indicate that an underestimation by about 10% may be due to biases in ocean carbon transport and

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3249 mixing from the surface mixed layer to the ocean interior (Goris et al., 2018, Terhaar et al., 2021, Bourgeois et  
 3250 al., 2022, Terhaar et al., 2022), biases in the chemical buffer capacity (Revelle factor) of the ocean (Vaittinada  
 3251 Ayar et al., 2022; Terhaar et al., 2022) and partly due to a late starting date of the simulations (mirrored in  
 3252 atmospheric CO<sub>2</sub> chosen for the preindustrial control simulation, Table S2, Bronselaer et al., 2017, Terhaar et  
 3253 al., 2022; 2024). Interestingly, and in contrast to the uncertainties in the surface CO<sub>2</sub> flux, we find the largest  
 3254 mismatch in interior ocean carbon accumulation in the tropics, with smaller contributions from the north, and the  
 3255 south. The large discrepancy in accumulation in the tropics highlights the role of interior ocean carbon  
 3256 redistribution for those inventories (Khatriwala et al., 2009, DeVries et al., 2023).

3257 The evaluation of the ocean estimates with the fCO<sub>2</sub> observations from the SOCAT y2024 dataset for the period  
 3258 1990-2023 shows an RMSE from annually detrended data of 0.2 to 2.4 μatm for the eight fCO<sub>2</sub>-products over  
 3259 the globe (Figure S2). The GOBMs RMSEs are larger and range from 2.7 to 4.9 μatm. The RMSEs are  
 3260 generally larger at high latitudes compared to the tropics, for both the fCO<sub>2</sub>-products and the GOBMs. The  
 3261 fCO<sub>2</sub>-products have RMSEs of 0.3 to 2.9 μatm in the Tropics, 0.6 to 2.4 μatm in the North, and 0.8 to 2.4 μatm  
 3262 in the South. Note that the fCO<sub>2</sub>-products are based on the SOCAT y2024 database, hence SOCAT is not an  
 3263 independent dataset for the evaluation of the fCO<sub>2</sub>-products. The GOBMs RMSEs are more spread across  
 3264 regions, ranging from 2.4 to 3.9 μatm in the tropics, 2.8 to 5.9 μatm in the North, and 2.7 to 6.0 μatm in the  
 3265 South. The higher RMSEs occur in regions with stronger climate variability, such as the northern and southern  
 3266 high latitudes (poleward of the subtropical gyres). Additionally, this year we evaluate the trends derived from a  
 3267 subset of fCO<sub>2</sub>-products by subsampling four GOBMs used in Friedlingstein et al. (2023; covering the period  
 3268 up to the year 2022) following the approach of Hauck et al. (2023a) and evaluating the air-sea CO<sub>2</sub> flux trend for  
 3269 the 2001-2021 period, i.e. the period of strong divergence in the air-sea CO<sub>2</sub> exchange excluding the final year  
 3270 to remove the tail effect, against trend biases identified by the GOBM reconstruction. The results indicate a  
 3271 relationship between reconstruction bias and strength of the decadal trends (see Figure S3), indicating a  
 3272 tendency of the fCO<sub>2</sub>-products ensemble to overestimate the air-sea CO<sub>2</sub> flux trends in agreement with a recent  
 3273 study by Mayot et al. (2024).

### 3274 3.7 Land Sink

#### 3275 3.7.1 Historical period 1850-2023

3276 Cumulated since 1850, the terrestrial CO<sub>2</sub> sink amounts to 220 ± 60 GtC, 31% of total anthropogenic emissions.  
 3277 As for the ocean, more than two thirds of this amount (150 ± 40 GtC) have been taken up by terrestrial  
 3278 ecosystems since 1960. Over the historical period, the sink increased in pace with the anthropogenic emissions  
 3279 increase (Figure 3).

#### 3280 3.7.2 Recent period 1960-2023

3281 The terrestrial CO<sub>2</sub> sink S<sub>LAND</sub> increased from 1.2 ± 0.5 GtC yr<sup>-1</sup> in the 1960s to 3.2 ± 0.9 GtC yr<sup>-1</sup> during 2014-  
 3282 2023, with important interannual variations of up to 2 GtC yr<sup>-1</sup> generally showing a decreased land sink during  
 3283 El Niño events (Figure 9), responsible for the corresponding enhanced growth rate in atmospheric CO<sub>2</sub>

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concentration. The larger land CO<sub>2</sub> sink during [2014-2023](#) compared to the 1960s is reproduced by all the DGVMs in response to the increase in both atmospheric CO<sub>2</sub>, nitrogen deposition, and the changes in climate, and is consistent with [the residual estimated](#) from the other budget terms ([E<sub>FOS</sub>+E<sub>LUC</sub>-G<sub>ATM</sub>-S<sub>OCEAN</sub>](#), Table 5).

Over the period 1960 to present the increase in the global terrestrial CO<sub>2</sub> sink is largely attributed to the CO<sub>2</sub> fertilisation effect (Prentice et al., 2001, Piao et al., 2009, Schimel et al., 2015) and increased nitrogen deposition (Huntzinger et al., 2017, O'Sullivan et al., 2019), directly stimulating plant photosynthesis and increased plant water use in water limited systems, with a small negative contribution of climate change (Figure [12](#)). There is a range of evidence to support a positive terrestrial carbon sink in response to increasing atmospheric CO<sub>2</sub>, albeit with uncertain magnitude (Walker et al., 2021). As expected from theory, the greatest CO<sub>2</sub> effect is simulated in the tropical forest regions, associated with warm temperatures and long growing seasons (Hickler et al., 2008) (Figure [12a](#)). However, evidence from tropical intact forest plots indicate an overall decline in the land sink across Amazonia (1985-2011), attributed to enhanced mortality offsetting productivity gains (Brienen et al., 2015, Hubau et al., 2020). During [2014-2023](#) the land sink is positive in all regions (Figure 6) with the exception of eastern Brazil, Bolivia, northern Venezuela, Southwest USA, central Europe and Central Asia, North and South Africa, and eastern Australia, where the negative effects of climate variability and change (i.e. reduced rainfall and/or increased temperature) counterbalance CO<sub>2</sub> effects. This is clearly visible on Figure [12](#) where the effects of CO<sub>2</sub> (Figure [12a](#)) and climate (Figure [12b](#)) as simulated by the DGVMs are isolated. The negative effect of climate [can be seen across the globe, and is particularly strong](#) in most of South America, Central America, Southwest US, Central Europe, western Sahel, southern Africa, Southeast Asia and southern China, and eastern Australia (Figure [12b](#)). Globally, over the [2014-2023](#) period, climate change reduces the land sink by  $0.87 \pm 0.56$  GtC yr<sup>-1</sup> (27% of S<sub>LAND</sub>).

Most DGVMs have similar S<sub>LAND</sub> averaged over [2014-2023](#), and 14/20 models fall within the 1σ range of the residual land sink [[1.8-3.7](#) GtC yr<sup>-1</sup>] (see Table 5), and all [models](#) but one are within the 2σ range [[0.8-4.6](#) GtC yr<sup>-1</sup>]. The ED model is an outlier, with a land sink estimate of [5.1](#) GtC yr<sup>-1</sup> [for the 2014-2023 period](#), driven by a strong CO<sub>2</sub> fertilisation effect ([6.3](#) GtC yr<sup>-1</sup> in the CO<sub>2</sub> only (S1) simulation). There are no direct global observations of the land sink (S<sub>LAND</sub>), or the CO<sub>2</sub> fertilisation effect, and so we are not yet in a position to rule out models based on component fluxes if [their](#) net land sink (S<sub>LAND</sub>-E<sub>LUC</sub>) is within the observational uncertainty provided by atmospheric [inversions or O<sub>2</sub> measurements](#) (Table 5). Furthermore, DGVMs were compared against a [model-data fusion based analysis](#) of the land carbon cycle (CARDAMOM) (Bloom and Williams, 2015; Bloom et al., 2016). Results suggest good correspondence between approaches at the interannual timescales, but divergence in the recent trend [in S<sub>LAND</sub>](#) with CARDAMOM simulating a stronger trend than the [DGVM multi-model mean](#) (Figure [9](#)).

Since 2020 the globe has experienced La Niña conditions which would be expected to lead to an increased land carbon sink. [This 3-year long period of La Niña conditions came to an end by the second half of 2023 and transitioned to an El Niño which lasted until mid-2024](#). A clear [transition from maximum to a minimum](#) in the global land sink is [evident in S<sub>LAND</sub>, from 2022 to 2023](#) and we find that a [El Niño- driven decrease](#) in tropical land sink is offset by a [smaller increase in the high latitude land sink](#). In the past years several regions

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experienced record-setting fire events (see also section 3.8.3). While global burned area has declined over the past decades mostly due to declining fire activity in savannas (Andela et al., 2017), forest fire emissions are rising and have the potential to counter the negative fire trend in savannas (Zheng et al., 2021). Noteworthy extreme fire events include the 2019-2020 Black Summer event in Australia (emissions of roughly 0.2 GtC; van der Velde et al., 2021), Siberia in 2021, where emissions approached 0.4 GtC or three times the 1997-2020 average according to GFED4s, and Canada in 2023 (Byrne et al., 2024). While other regions, including Western US and Mediterranean Europe, also experienced intense fire seasons in 2021 their emissions are substantially lower.

Despite these regional negative effects of climate change on  $S_{LAND}$ , the efficiency of land to remove anthropogenic  $CO_2$  emissions has remained broadly constant over the last six decades, with a land-borne fraction ( $S_{LAND}/(E_{FOS}+E_{LUC})$ ) of around 30% (Figure 10b).

### 3.7.3 Final year 2023

The terrestrial  $CO_2$  sink from the DGVMs ensemble  $S_{LAND}$  was  $2.3 \pm 1.0$  GtC in 2023, 41% below the 2022 La Niña induced strong sink of  $3.9 \pm 1.0$  GtC, and also below the 2014-2023 average of  $3.2 \pm 0.9$  GtC  $yr^{-1}$  (Figure 4, Table 7). We estimate that the 2023 land sink was the lowest since 2015. The severe reduction in the land sink in 2023 is likely driven by the El Niño conditions, leading to a 58% reduction in  $S_{LAND}$  in the tropics (30N-30S) from 2.8 GtC in 2022 to 1.2 GtC in 2023. This is combined with intense wildfires in Canada that led to a significant  $CO_2$  source (see also Section 3.8.3). We note that the  $S_{LAND}$  DGVMs estimate for 2023 of  $2.3 \pm 1.0$  GtC is very similar to the  $2.2 \pm 1.0$  GtC  $yr^{-1}$  estimate from the residual sink from the global budget ( $E_{FOS}+E_{LUC}-G_{ATM-SOCEAN}$ , Table 5).

### 3.7.4 Year 2024 Projection

Using a feed-forward neural network method we project a land sink of 3.2 GtC for 2024, 0.9 GtC larger than the 2023 estimate. As for the ocean sink, we attribute this to the transition from the El Niño conditions in 2023 to a neutral state. The ESMs do not provide an additional estimate of  $S_{LAND}$  as they only simulate the net atmosphere-land carbon flux ( $S_{LAND}-E_{LUC}$ ).

### 3.7.5 Land Models Evaluation

The evaluation of the DGVMs shows generally higher agreement across models for runoff, and to a lesser extent for GPP, and ecosystem respiration. These conclusions are supported by a more comprehensive analysis of DGVM performance in comparison with benchmark data (Sitch et al., 2024). A relative comparison of DGVM performance (Figure S4) suggests several DGVMs (CABLE-POP, CLASSIC, OCN, ORCHIDEE) may outperform others at multiple carbon and water cycle benchmarks. However, results from Seiler et al., 2022, also show how DGVM differences are often of similar magnitude compared with the range across observational datasets. All models score high enough over the metrics tests to support their use here. There are a few

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3474 anomalously low scores for individual metrics from a single model, and these can direct the effort to improve  
3475 models for use in future budgets.

### 3476 3.8 Partitioning the carbon sinks

#### 3477 3.8.1 Global sinks and spread of estimates

3478 In the period [2014-2023](#), the bottom-up view of global net ocean and land carbon sinks provided by the GCB,  
3479 SOCEAN for the ocean and SLAND-ELUC for the land, agrees closely with the top-down global carbon sinks  
3480 delivered by the atmospheric inversions. This is shown in Figure [13](#), which visualises the individual decadal  
3481 mean atmosphere-land and atmosphere-ocean fluxes from each, along with the constraints on their sum offered  
3482 by the global fossil CO<sub>2</sub> emissions flux minus the atmospheric growth rate ( $E_{FOS} - G_{ATM}$ ,  $4.4 \pm 0.5$  Gt C yr<sup>-1</sup>,  
3483 Table 7, shown as diagonal line on Figure [13](#)). The GCB estimate for net atmosphere-to-surface flux (SOCEAN +  
3484 SLAND - ELUC) during [2014-2023](#) is  $4.9 \pm 1.2$  Gt C yr<sup>-1</sup> (Table 7), with the difference to the diagonal representing  
3485 the budget imbalance (BIM) of 0.4 Gt C yr<sup>-1</sup> discussed in Section 3.9. By virtue of the inversion methodology, the  
3486 [atmospheric inversions estimate of the net atmosphere-to-surface flux during 2014-2023 is 4.5 Gt C yr<sup>-1</sup>, with a](#)  
3487 [< 0.1 Gt C yr<sup>-1</sup> imbalance, and thus scatter across the diagonal, with](#) inverse models trading land for ocean fluxes  
3488 in their solution. The independent constraint on the net atmosphere-to-surface flux based on atmospheric O<sub>2</sub> [by](#)  
3489 [design also closes the balance and](#) is  $4.5 \pm 0.9$  Gt C yr<sup>-1</sup> over the [2014-2023](#) period (orange symbol on Figure  
3490 [13](#)), while the ESMs estimate for the net atmosphere-to-surface flux over that period [average to 4.7 \[3.0, 5.8\]](#)  
3491 [GtC yr<sup>-1</sup>](#) (Tables 5 and 6).

3492 The distributions based on the individual models and [fCO<sub>2</sub>](#)-products reveal substantial spread but converge near  
3493 the decadal means quoted in Tables 5 to 7. Sink estimates for SOCEAN and from inverse systems are mostly non-  
3494 Gaussian, while the ensemble of DGVMs appears more normally distributed justifying the use of a multi-model  
3495 mean and standard deviation for their errors in the budget. Noteworthy is that the tails of the distributions  
3496 provided by the land and ocean bottom-up estimates would not agree with the global constraint provided by the  
3497 fossil fuel emissions and the observed atmospheric CO<sub>2</sub> growth rate. This illustrates the power of the  
3498 atmospheric joint constraint from G<sub>ATM</sub> and the global CO<sub>2</sub> observation network it derives from.

##### 3499 3.8.1.1 Net atmosphere-to-land flux

3500 The GCB [estimate of the](#) net atmosphere-to-land [flux](#) (SLAND - ELUC), calculated as the difference between  
3501 SLAND from the DGVMs and ELUC from the bookkeeping models, amounts to a  $2.1 \pm 1.1$  GtC yr<sup>-1</sup> sink during  
3502 [2014-2023](#) (Table 5). Estimates of net atmosphere-to-land [flux](#) (SLAND - ELUC) from the DGVMs alone ( $1.7 \pm$   
3503  $0.6$  GtC yr<sup>-1</sup>, Table 5, green symbol on Figure [13](#)) are slightly lower, [although](#) within the uncertainty of the GCB  
3504 estimate and also [within uncertainty of](#) the global carbon budget constraint ( $E_{FOS} - G_{ATM} - \text{SOCEAN}$ ,  $1.6 \pm 0.6$  GtC  
3505 yr<sup>-1</sup>; Table 7). [Also, for 2014-2023, the inversions estimate the net atmosphere-to-land flux is a 1.4 \[0.3, 2.2\]](#)  
3506 [GtC yr<sup>-1</sup>, sink, slightly lower than the mean of the DGVMs estimates \(purple versus grey symbols on Figure 13\).](#)  
3507 [The independent constraint based on atmospheric O<sub>2</sub> is even lower, 1.0 ± 0.8 GtC yr<sup>-1</sup> \(orange symbol in Figure](#)  
3508 [13\), although its large uncertainty overlaps with the uncertainty range from other approaches. Last, the ESMs](#)

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estimate for the net atmosphere-to-land flux during 2014-2023 is a  $2.2 [0.3, 3.6]$  GtC yr<sup>-1</sup> sink, more consistent with the GCB estimates of  $S_{\text{LAND}} - E_{\text{LUC}}$  (Figure 14 top row).

As discussed in Section 3.5.3, the atmospheric growth rate of CO<sub>2</sub> was very high in 2023, 5.9 GtC (2.79 ppm) the 4<sup>th</sup> largest on record. Both DGVMs and inversions assign this large CO<sub>2</sub> growth rate to a severe decrease of the net atmosphere to land flux, and in particular in the tropics (Figure 14). DGVMs simulate a 2023 global the net atmosphere-to-land flux of  $1 \downarrow$  GtC yr<sup>-1</sup>, a 55% decline relative to the 2.4 GtC yr<sup>-1</sup> sink in 2022, primarily driven by the severe reduction in  $S_{\text{LAND}}$  (-41%, see Section 3.7.3). The tropics (30N-30S) are recording a dramatic decrease in the net atmosphere-to-land flux from 1.5 GtC yr<sup>-1</sup> in 2022 to 0.1 GtC yr<sup>-1</sup> in 2023. The atmospheric inversion shows a similar story with the global net atmosphere-to-land flux declining from 2.6 GtC yr<sup>-1</sup> in 2022 to 0.9 GtC yr<sup>-1</sup> in 2023 (-64%), with the tropics turning from a 1.0 GtC yr<sup>-1</sup> sink in 2022 to a 0.4 GtC yr<sup>-1</sup> source in 2023. Our results are broadly consistent with the Ke et al. (2024) study which reported a global atmosphere-to-land flux of  $0.4 \pm 0.2$  GtC yr<sup>-1</sup> in 2023.

In addition to the large decline of the tropical land uptake, the northern extra tropics experienced warmer than average conditions, in particular in the summer over North America and North Eurasia. In Canada alone, 2023 led to enhanced CO<sub>2</sub> release due to fires of 0.5-0.8 GtC yr<sup>-1</sup> (see Section 3.8.3). The atmospheric inversions do simulate a slight reduction of the atmosphere-to-land flux in the northern extra-tropics (north of 30°N), from 1.6 GtC yr<sup>-1</sup> in 2022 to 1.4 GtC yr<sup>-1</sup> in 2023, while the DGVM fail to capture this pattern, with a simulated northern extra-tropics net atmosphere-to-land flux larger in 2023 than in 2022 (1.0 vs 0.7 GtC yr<sup>-1</sup>).

### 3.8.1.2 Net atmosphere-to-ocean flux

For the 2014-2023 period, the GOBMs ( $2.6 \pm 0.4$  GtC yr<sup>-1</sup>) produce a lower estimate for  $S_{\text{OCEAN}}$  than the  $f\text{CO}_2$ -products with  $3.1 [2.9, 3.7]$  GtC yr<sup>-1</sup>, which shows up in Figure 13 as separate peaks in the distribution from the GOBMs (dark blue symbols) and from the  $f\text{CO}_2$ -products (light blue symbols). Atmospheric inversions ( $3 \downarrow [2.4, 4.1]$  GtC yr<sup>-1</sup>) suggest an ocean uptake more in line with the  $f\text{CO}_2$ -products for the recent decade (Table 7), although the inversions range includes both the GOBMs and  $f\text{CO}_2$ -products estimates (Figure 14 top row) and the inversions are not fully independent as 6 out of 10 inversions covering the last decade use  $f\text{CO}_2$ -products as ocean priors and one uses a GOBM (Table S4). The independent constraint based on atmospheric O<sub>2</sub> ( $3.4 \pm 0.5$  GtC yr<sup>-1</sup>) is at the high end of the distribution of the other methods. However, as mentioned in section 2.8, the O<sub>2</sub> method requires a correction for global air-sea O<sub>2</sub> flux, which induces a non-negligible uncertainty on the decadal estimates (about 0.5 GtC yr<sup>-1</sup>). The large growth in the ocean carbon sink from O<sub>2</sub> is compatible with the GOBMs and  $f\text{CO}_2$ -products estimates when accounting for their uncertainty ranges. Lastly, the ESMs estimate,  $2.5 [2.2, 2.8]$  GtC yr<sup>-1</sup>, suggest a moderate ocean carbon sink, comparable to the GOBMs estimate with regard to mean and spread. We caution that the riverine transport of carbon taken up on land and outgassing from the ocean, accounted for here, is a substantial ( $0.65 \pm 0.3$  GtC yr<sup>-1</sup>) and uncertain term (Crisp et al., 2022; Gruber et al., 2023; DeVries et al., 2023) that separates the GOBMs, ESMs and oxygen-based estimates on the one hand from the  $f\text{CO}_2$ -products and atmospheric inversions on the other hand.

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### 3607 3.8.2 Regional partitioning

3608 Figure 14 shows the latitudinal partitioning of the global atmosphere-to-ocean ( $S_{\text{OCEAN}}$ ), atmosphere-to-land  
3609 ( $S_{\text{LAND}} - E_{\text{LUC}}$ ), and their sum ( $S_{\text{OCEAN}} + S_{\text{LAND}} - E_{\text{LUC}}$ ) according to the estimates from GOBMs and ocean  
3610  $f\text{CO}_2$ -products ( $S_{\text{OCEAN}}$ ), DGVMs ( $S_{\text{LAND}} - E_{\text{LUC}}$ ), and from atmospheric inversions ( $S_{\text{OCEAN}}$  and  $S_{\text{LAND}} - E_{\text{LUC}}$ ).

#### 3611 3.8.2.1 North

3612 Despite being one of the most densely observed and studied regions of our globe, annual mean carbon sink  
3613 estimates in the northern extra-tropics (north of  $30^\circ\text{N}$ ) continue to differ. The atmospheric inversions suggest an  
3614 atmosphere-to-surface sink ( $S_{\text{OCEAN}} + S_{\text{LAND}} - E_{\text{LUC}}$ ) for 2014–2023 of  $2.6 [2.0 \text{ to } 3.4] \text{ GtC yr}^{-1}$ , which is slightly  
3615 higher than the process models' estimate of  $2.2 \pm 0.4 \text{ GtC yr}^{-1}$  (Figure 14). The GOBMs ( $1.2 \pm 0.2 \text{ GtC yr}^{-1}$ ),  
3616  $f\text{CO}_2$ -products ( $1.4 [1.3 \text{ to } 1.5] \text{ GtC yr}^{-1}$ ), and inversion systems ( $1.2 [0.9 \text{ to } 1.4] \text{ GtC yr}^{-1}$ ) produce largely  
3617 consistent estimates of the ocean sink. However, the larger flux in the  $f\text{CO}_2$ -products may be related to data  
3618 sparsity (Mayot et al., 2024). Thus, the difference mainly arises from the net land flux ( $S_{\text{LAND}} - E_{\text{LUC}}$ ) estimate,  
3619 which is  $1.0 \pm 0.4 \text{ GtC yr}^{-1}$  in the DGVMs compared to  $1.5 [0.6 \text{ to } 2.3] \text{ GtC yr}^{-1}$  in the atmospheric inversions  
3620 (Figure 14, second row).

3621 Discrepancies in the northern land fluxes conforms with persistent issues surrounding the quantification of the  
3622 drivers of the global net land  $\text{CO}_2$  flux (Armeth et al., 2017; Huntzinger et al., 2017; O'Sullivan et al., 2022) and  
3623 the distribution of atmosphere-to-land fluxes between the tropics and high northern latitudes (Baccini et al.,  
3624 2017; Schimel et al., 2015; Stephens et al., 2007; Ciais et al., 2019; Gaubert et al., 2019).

3625 In the northern extra-tropics, the process models, inversions, and  $f\text{CO}_2$ -products consistently suggest that most  
3626 of the interannual variability stems from the land (Figure 14). Inversions generally agree on the magnitude of  
3627 interannual variations (IAV) over land, more so than DGVMs ( $0.29\text{--}0.32$  vs  $0.14\text{--}0.63 \text{ GtC yr}^{-1}$ , averaged over  
3628 1990–2023).

#### 3629 3.8.2.2 Tropics

3630 In the tropics ( $30^\circ\text{S}$ – $30^\circ\text{N}$ ), both the atmospheric inversions and process models estimate a net carbon balance  
3631 ( $S_{\text{OCEAN}} + S_{\text{LAND}} - E_{\text{LUC}}$ ) that is relatively close to neutral over the past decade (inversions:  $0.3 [-0.4, 0.9] \text{ GtC yr}^{-1}$ ,  
3632 process models:  $0.6 \pm 0.6 \text{ GtC yr}^{-1}$ ). The GOBMs ( $-0.03 \pm 0.3 \text{ GtC yr}^{-1}$ ),  $f\text{CO}_2$ -products ( $0.3 [0.1, 0.6] \text{ GtC yr}^{-1}$ ),  
3633 and inversion systems ( $0.3 [-0.1, 0.8] \text{ GtC yr}^{-1}$ ) indicate a neutral to positive tropical ocean flux (see Figure  
3634 S1 for spatial patterns). DGVMs indicate a net land sink ( $S_{\text{LAND}} - E_{\text{LUC}}$ ) of  $0.6 \pm 0.4 \text{ GtC yr}^{-1}$ , whereas the  
3635 inversion systems indicate a neutral net land flux, although with large model spread ( $-0.0 [-0.9, 0.8] \text{ GtC yr}^{-1}$ ,  
3636 Figure 14, third row).

3637 The tropical lands are the origin of most of the atmospheric  $\text{CO}_2$  interannual variability (Ahlström et al., 2015),  
3638 consistently among the process models and inversions (Figure 14). The interannual variability in the tropics is  
3639 similar among the ocean  $f\text{CO}_2$ -products ( $0.06\text{--}0.16 \text{ GtC yr}^{-1}$ ) and the GOBMs ( $0.07\text{--}0.16 \text{ GtC yr}^{-1}$ , Figure S2).  
3640 The DGVMs and inversions indicate that atmosphere-to-land  $\text{CO}_2$  fluxes are more variable than atmosphere-to-

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[76]

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3724 ocean CO<sub>2</sub> fluxes in the tropics, with interannual variability of 0.37 to 1.33 and 0.86-0.96 GtC yr<sup>-1</sup> for DGVMs  
3725 and inversions, respectively.

### 3726 3.8.2.3 South

3727 In the southern extra-tropics (south of 30°S), the atmospheric inversions suggest a net atmosphere-to-surface  
3728 sink (SOCEAN+SLAND-ELUC) for 2014-2023 of 1.5 [1.2, 1.9] GtC yr<sup>-1</sup>, identical to the process models' estimate of  
3729 1.5 ± 0.4 GtC yr<sup>-1</sup> (Figure 14). An approximately neutral net land flux (SLAND-ELUC) for the southern extra-  
3730 tropics is estimated by both the DGVMs (0.05 ± 0.1 GtC yr<sup>-1</sup>) and the inversion systems (-0.03 [-0.11, 0.08] GtC  
3731 yr<sup>-1</sup>). This means nearly all carbon uptake is due to oceanic sinks south of 30°S. The Southern Ocean flux in the  
3732 fCO<sub>2</sub>-products (1.5 [1.3, 1.7 GtC] yr<sup>-1</sup>) and inversion estimates (1.6 [1.2, 1.9] GtCyr-1) is marginally higher than  
3733 in the GOBMs (1.4 ± 0.4 GtC yr<sup>-1</sup>) (Figure 14, bottom row). This agreement is subject to the choice of the river  
3734 flux adjustment (Lacroix et al., 2020, Hauck et al., 2023b). Nevertheless, the time-series of atmospheric  
3735 inversions and fCO<sub>2</sub>-products diverge from the GOBMs. A substantial overestimation of the trends in the fCO<sub>2</sub>-  
3736 products could be explained by sparse and unevenly distributed observations, especially in wintertime (Figure  
3737 S1; Hauck et al., 2023a; Gloege et al., 2021). Model biases may contribute as well, with biases in mode water  
3738 formation, stratification, and the chemical buffer capacity known to play a role in Earth System Models (Terhaar  
3739 et al., 2021, Bourgeois et al., 2022, Terhaar et al., 2022).

3740 The interannual variability in the southern extra-tropics is low because of the dominance of ocean areas with  
3741 low variability compared to land areas. The split between land (SLAND-ELUC) and ocean (SOCEAN) shows a  
3742 substantial contribution to variability in the south coming from the land, with no consistency between the  
3743 DGVMs and the inversions or among inversions. This is expected due to the difficulty of separating exactly the  
3744 land and oceanic fluxes when viewed from atmospheric observations alone. The SOCEAN interannual variability  
3745 was found to be higher in the fCO<sub>2</sub>-products (0.04-0.20 GtC yr<sup>-1</sup>) compared to GOBMs (0.04 to 0.06 GtC yr<sup>-1</sup>)  
3746 in 1990-2023 (Figure S2). Inversions give an interannual variability of 0.10 to 0.13 GtC yr<sup>-1</sup>. Model  
3747 subsampling experiments recently illustrated that fCO<sub>2</sub>-products may overestimate decadal variability in the  
3748 Southern Ocean carbon sink by 30% and the trend since 2000 by 50-130% due to data sparsity, based on one  
3749 and two fCO<sub>2</sub>-products with strong variability (Gloege et al., 2021, Hauck et al., 2023a). The trend benchmark  
3750 test using the method of Hauck et al., (2023a) and a subset of 6 fCO<sub>2</sub>-products confirms the sensitivity of the  
3751 decadal trends in fCO<sub>2</sub>-products to reconstruction biases, particularly in the Southern Ocean, indicating an  
3752 overestimation of the ensemble mean trend. However, we also find compensating positive biases in the  
3753 ensemble so that the ensemble mean bias is smaller than the bias from some individual fCO<sub>2</sub>-products.

### 3754 3.8.2.4 RECCAP2 regions

3755 Aligning with the RECCAP-2 initiative (Ciais et al., 2022; Poulter et al., 2022; DeVries et al., 2023), we  
3756 provide a breakdown of this GCB paper estimate of the ELUC, SLAND, Net land (SLAND - ELUC), and SOCEAN fluxes  
3757 over the 10 land, and 5 ocean RECCAP-2 regions, averaged over the period 2014-2023 (Figure 15). The  
3758 DGVMs and inversions suggest a positive net land sink in all regions, except for South America and Africa,

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where the inversions indicate a small net source of respectively  $-0.1 [-0.8, 0.3]$   $\text{GtC yr}^{-1}$  and  $-0.3 [-0.7, -0.1]$   $\text{GtC yr}^{-1}$ , compared to a small sink of  $0.1 \pm 0.3$   $\text{GtC yr}^{-1}$  and  $0.3 \pm 0.1$   $\text{GtC yr}^{-1}$  for the DGVMs. However, for South America, there is substantial uncertainty in both products (ensembles span zero). For the DGVMs, this is driven by uncertainty in both  $S_{\text{LAND}}$  ( $0.5 \pm 0.4$   $\text{GtC yr}^{-1}$ ) and  $E_{\text{LUC}}$  ( $0.4 \pm 0.2$   $\text{GtC yr}^{-1}$ ). The bookkeeping models also suggest an  $E_{\text{LUC}}$  source of around  $0.4$   $\text{GtC yr}^{-1}$  in South America and Africa, in line with the DGVMs estimates. Bookkeeping models and DGVMs similarly estimate a source of  $0.3-0.4$   $\text{GtC yr}^{-1}$  in Southeast Asia, with DGVMs suggesting a *small* net land sink ( $0.1 \pm 0.1$   $\text{GtC yr}^{-1}$ ). This *is similar to* the inversion *mean* estimate of a  $0.1 [-0.3, 0.8]$   $\text{GtC yr}^{-1}$  sink, although the *inversion* spread is substantial. The inversions suggest the largest net land sinks are located in North America ( $0.5 [-0.1, 1.0]$   $\text{GtC yr}^{-1}$ ), Russia ( $0.6 [0.1, 0.9]$   $\text{GtC yr}^{-1}$ ), and East Asia ( $0.4 [-0.2, 1.3]$   $\text{GtC yr}^{-1}$ ). This agrees well with the DGVMs in North America ( $0.4 \pm 0.1$   $\text{GtC yr}^{-1}$ ), which indicate a large natural land sink ( $S_{\text{LAND}}$ ) of  $0.6 \pm 0.2$   $\text{GtC yr}^{-1}$ , being slightly reduced by land-use related carbon losses ( $0.2 \pm 0.1$   $\text{GtC yr}^{-1}$ ). The DGVMs suggest a smaller net land sink in Russia compared to inversions ( $0.3 \pm 0.2$   $\text{GtC yr}^{-1}$ ), and a similar net sink in East Asia ( $0.2 \pm 0.1$   $\text{GtC yr}^{-1}$ ).

There is generally a higher level of agreement in the estimates of regional  $S_{\text{OCEAN}}$  between the different data streams (GOBMs,  $f\text{CO}_2$ -products and atmospheric inversions) on decadal scale, compared to the agreement between the different land flux estimates. All data streams agree that the largest contribution to  $S_{\text{OCEAN}}$  stems from the Southern Ocean due to a combination of high flux density and large surface area, but with important contributions also from the Atlantic (high flux density) and Pacific (large area) basins. In the Southern Ocean, GOBMs suggest a sink of  $1.0 \pm 0.3$   $\text{GtC yr}^{-1}$ , in line with the  $f\text{CO}_2$ -products ( $1.0 [0.8, 1.3]$   $\text{GtC yr}^{-1}$ ) and atmospheric inversions ( $1.0 [0.7, 1.4]$   $\text{GtC yr}^{-1}$ ). There is similar agreement in the Pacific ocean, with GOBMs,  $f\text{CO}_2$ -products, and atmospheric inversions indicating a sink of  $0.6 \pm 0.2$   $\text{GtC yr}^{-1}$ ,  $0.7 [0.6, 1.0]$   $\text{GtC yr}^{-1}$ , and  $0.6 [0.1, 1.0]$   $\text{GtC yr}^{-1}$ , respectively. However, in the Atlantic ocean, GOBMs simulate a sink of  $0.5 \pm 0.1$   $\text{GtC yr}^{-1}$ , noticeably lower than both the  $f\text{CO}_2$ -products ( $0.8 [0.7, 1.0]$   $\text{GtC yr}^{-1}$ ) and atmospheric inversions ( $0.7 [0.4, 1.1]$   $\text{GtC yr}^{-1}$ ). It is important to note the  $f\text{CO}_2$ -products and atmospheric inversions have a substantial and uncertain river flux adjustment in the Atlantic ocean ( $0.3$   $\text{GtC yr}^{-1}$ ) that also leads to a mean offset between GOBMs and  $f\text{CO}_2$ -products/inversions in the latitude band of the tropics (Figure 14). The Indian Ocean due its smaller size and the Arctic Ocean due to its size and sea-ice cover that prevents air-sea gas-exchange are responsible for smaller but non negligible  $S_{\text{OCEAN}}$  fluxes (Indian Ocean:  $0.3 [0.2, 0.3]$   $\text{GtC yr}^{-1}$ ,  $0.3 [0.3, 0.4]$   $\text{GtC yr}^{-1}$ , and  $0.4 [0.3, 0.6]$   $\text{GtC yr}^{-1}$  for GOBMs,  $f\text{CO}_2$ -products, and atmospheric inversions, respectively, and Arctic Ocean:  $0.1 [0.1, 0.1]$   $\text{GtC yr}^{-1}$ ,  $0.2 [0.1, 0.2]$   $\text{GtC yr}^{-1}$ , and  $0.1 [0.1, 0.2]$   $\text{GtC yr}^{-1}$  for GOBMs,  $f\text{CO}_2$ -products, and atmospheric inversions, respectively). Note that the  $S_{\text{OCEAN}}$  numbers presented here deviate from numbers reported in RECCAP-2 where the net air-sea  $\text{CO}_2$  flux is reported (i.e. without river flux adjustment for  $f\text{CO}_2$ -products and inversions, and with river flux adjustment subtracted from GOBMs in most chapters, or comparing unadjusted *datasets* with discussion of uncertain regional riverine fluxes as major uncertainty, e.g. Sarma et al., 2023, DeVries et al., 2023).

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### 3.8.2.5 Tropical vs northern land uptake

A continuing conundrum is the partitioning of the global atmosphere-land flux between the northern hemisphere land and the tropical land (Stephens et al., 2017; Pan et al., 2011; Gaubert et al., 2019). It is of importance because each region has its own history of land-use change, climate drivers, and impact of increasing atmospheric CO<sub>2</sub> and nitrogen deposition. Quantifying the magnitude of each sink is a prerequisite to understanding how each individual driver impacts the tropical and mid/high-latitude carbon balance.

We define the North-South (N-S) difference as net atmosphere-land flux north of 30°N minus the net atmosphere-land flux south of 30°N. For the inversions, the N-S difference is 1.50 [0.05,3.0] GtC yr<sup>-1</sup> across this year's inversion ensemble. An apparent clustering of six satellite-driven solutions towards a common NH land sink noted in GCB2023 is no longer clear.

In the ensemble of DGVMs the N-S difference is  $0.4 \pm 0.5$  GtC yr<sup>-1</sup>, a much narrower range than the one from atmospheric inversions. Only three out of twenty DGVMs have a N-S difference larger than 1.0 GtC yr<sup>-1</sup>, compared to half of the inversion systems simulating a difference at least this large. The smaller spread across DGVMs than across inversions is to be expected as there is no correlation between Northern and Tropical land sinks in the DGVMs as opposed to the inversions where the sum of the two regions being well-constrained by atmospheric observations leads to an anti-correlation between these two regions. This atmospheric N-S gradient could be used as an additional way to evaluate tropical and NH uptake in DGVMs, if their fluxes were combined with multiple transport models. Vice versa, the much smaller spread in the N-S difference between the DGVMs could help to scrutinise the inverse systems further. For example, a large northern land sink and a tropical land source in an inversion would suggest a large sensitivity to CO<sub>2</sub> fertilisation (the dominant factor driving the land sinks) for Northern ecosystems, which would be not mirrored by tropical ecosystems. Such a combination could be hard to reconcile with the process understanding gained from the DGVM ensembles and independent measurements (e.g. Free Air CO<sub>2</sub> Enrichment experiments).

### 3.8.3 Fire Emissions in 2024

Fire emissions so far in 2024 have been above the average of recent decades, chiefly due to synchronous large emissions fluxes from North and South America. Figure S9 shows global and regional emissions estimates for the period 1st Jan-30th September in each year 2003-2024. Estimates derive from two global fire emissions products: the global fire emissions database (GFED, version 4.1s; van der Werf et al., 2017), and the global fire assimilation system (GFAS, operated by the Copernicus Atmosphere Service; Kaiser et al., 2012). The two products estimate that global emissions from fires were 1.6-2.2 GtC yr<sup>-1</sup> during January-September 2024. These estimates are 11-32% above the 2014-2023 average for the same months (1.5-1.7 GtC yr<sup>-1</sup>). In the GFED4.1s product, the year-to-date emissions in 2024 were highest since 2003, exceeding even the large emissions estimate of 2023, whereas the GFAS product showed lower emissions in 2024 than in 2023 and six other years since 2003.

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The pattern of high fire emissions from Canada in 2023, which were record-breaking (Jones et al., 2024b, Byrne et al., 2024), continued into 2024. In January-September 2024, emissions from Canada (0.2-0.3 GtC yr<sup>-1</sup>) were half as great as in the same months of 2023 (0.5-0.8 GtC yr<sup>-1</sup>) but still 2.1-2.3 times the average of January-September periods in 2014-2023 (and 4-6 times greater than the average of those months in 2003-2022 [excluding the record-breaking year in 2023]; Figure S9). The continued anomaly in Canada propagated to the northern hemisphere, where emissions of 0.5-0.6 GtC yr<sup>-1</sup> were 26-44% above the average of 2014-2023.

In January-September 2024, fire emissions from South America (0.4-0.6 GtC yr<sup>-1</sup>) were 94-164% above the average of January-September periods in 2014-2023, marking 2024 out as a year with synchronous high fire emissions across the Americas. Emissions from Brazil in January-September 2024 (0.2-0.3 GtC yr<sup>-1</sup>) were 91-118% above the average of January-September periods of 2014-2023 and were at a level not seen since the major drought year of 2010 (Figure S9; Aragão et al., 2018, Silva Junior et al., 2019). In 2023, deforestation fire activity in the Brazilian Amazon was below the average levels recorded in national recording systems and attributed to renewed environmental policy implementation, however the fall in Amazon deforestation fire activity was largely offset by above-average wildfires related to historic drought (Mataveli et al. 2024). According to the National Center for Monitoring and Early Warning of Natural Disasters (CEMADEN), drought conditions continued into 2024 and the current drought is the most intense and widespread Brazil has experienced since records began in 1950 (CEMADEN, 2024), prompting large wildfires anomalies across the Amazon, Cerrado and Pantanal regions (INPE, 2024).

Emissions anomalies in Africa strongly influence global totals because the continent typically contributed 41-47% of global fire emissions during 2014-2023 (average of January-September periods). GFAS suggests that fire emissions in Africa through September 2024 (0.6 GtC yr<sup>-1</sup>) were slightly below the average of 2014-2023, whereas GFED4.1s suggests that fire emissions through September 2024 were slightly above the average of 2014-2023 (0.8 GtC yr<sup>-1</sup>).

Tropical fire emissions through September 2024 (1.1-1.6 GtC yr<sup>-1</sup>) accounted for 69-74% of the global total emissions, which is close to the average of the 2014-2023 period (1.1-1.2 GtC yr<sup>-1</sup>; 72-75%). This marks a return to a more typical distribution of fire emissions between the tropics and extratropics after the tropical contribution fell to just 55-59% during January-September 2023 (Figure S9).

We caution that the fire emissions fluxes presented here should not be compared directly with other fluxes of the budget (e.g. S<sub>LAND</sub> or E<sub>LUC</sub>) due to incompatibilities between the observable fire emission fluxes and what is quantified in the S<sub>LAND</sub> and E<sub>LUC</sub> components of the budget. The fire emission estimates from global fire products relate to all fire types that can be observed in Earth Observations (Giglio et al., 2018; Randerson et al., 2012; Kaiser et al., 2012), including (i) fires occurring as part of natural disturbance-recovery cycles that would also have occurred in the pre-industrial period (Yue et al., 2016; Keeley and Pausas, 2019; Zou et al., 2019), (ii) fires occurring above and beyond natural disturbance-recovery cycle due to changes in climate, CO<sub>2</sub> and N fertilisation and to an increased frequency of extreme drought and heatwave events (Abatzoglou et al., 2019; Jones et al., 2022; Zheng et al., 2021; Burton et al., 2024), and (iii) fires occurring in relation to land use and land use change, such as deforestation fires and agricultural fires (van der Werf et al., 2010; Magi et al., 2012).

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In the context of the global carbon budget, only the portion of fire emissions associated with (ii) should be included in the  $S_{\text{LAND}}$  component, and fire emissions associated with (iii) should already be accounted for in the  $E_{\text{LUC}}$  component. Emissions associated with (i) should not be included in the global carbon budget. It is not currently possible to derive specific estimates for fluxes (i), (ii), and (iii) using global fire emission products such as GFED or GFAS. In addition, the fire emissions estimates from global fire emissions products represent a gross flux of carbon to the atmosphere, whereas the  $S_{\text{LAND}}$  component of the budget is a net flux that should also include post-fire recovery fluxes. Even if emissions from fires of type (ii) could be separated from those of type (i), these fluxes may be partially or wholly offset in subsequent years by post-fire fluxes as vegetation recovers, sequestering carbon from the atmosphere to the terrestrial biosphere (Yue et al., 2016; Jones et al., 2024c). Increases in forest fire emissions and severity (emissions per unit area) from globally during the past two decades have highlighted the increasing potential for fire emissions fluxes to outweigh post-fire recovery fluxes, though long-term monitoring of vegetation recovery is required to quantify the net effect on terrestrial C storage (Jones et al., 2024c).

### 3.9 Closing the Global Carbon Cycle

#### 3.9.1 Partitioning of Cumulative Emissions and Sink Fluxes

Emissions during the period 1850–2023 amounted to  $710 \pm 70$  GtC and were partitioned among the atmosphere ( $285 \pm 5$  GtC; 40%), ocean ( $185 \pm 35$  GtC; 26%), and land ( $220 \pm 60$  GtC; 32%). The cumulative land sink is almost equal to the cumulative land-use emissions ( $225 \pm 65$  GtC), making the global land nearly neutral over the whole 1850–2023 period (Figure 3).

The use of nearly independent estimates for the individual terms of the global carbon budget shows a cumulative budget imbalance of  $25$  GtC (3% of total emissions) during 1850–2023 (Figure 3, Table 8), which, if correct, suggests that emissions could be slightly too high by the same proportion or that the combined land and ocean sinks are slightly underestimated (by about 6%), although these are well within the uncertainty range of each component of the budget. Nevertheless, part of the imbalance could originate from the estimation of significant increase in  $E_{\text{FOS}}$  and  $E_{\text{LUC}}$  between the mid 1920s and the mid 1960s which is unmatched by a similar growth in atmospheric  $\text{CO}_2$  concentration as recorded in ice cores (Figure 3). However, the known loss of additional sink capacity of 30–40 GtC (over the 1850–2020 period) due to reduced forest cover has not been accounted for in our method and would exacerbate the budget imbalance (see Section 2.10 and Supplement S.6.4).

For the more recent 1960–2023 period where direct atmospheric  $\text{CO}_2$  measurements are available, total emissions ( $E_{\text{FOS}} + E_{\text{LUC}}$ ) amounted to  $500 \pm 50$  GtC, of which  $410 \pm 20$  GtC (82%) were caused by fossil  $\text{CO}_2$  emissions, and  $90 \pm 45$  GtC (18%) by land-use change (Table 8). The total emissions were partitioned among the atmosphere ( $220 \pm 5$  GtC; 45%), ocean ( $130 \pm 26$  GtC; 25%), and the land ( $150 \pm 40$  GtC; 30%), with a near zero ( $<1$  GtC) unattributed budget imbalance. All components except land-use change emissions have significantly grown since 1960, with important interannual variability in the growth rate in atmospheric  $\text{CO}_2$  concentration and in the land  $\text{CO}_2$  sink (Figure 4), and some decadal variability in all terms (Table 7). Differences with previous budget releases are documented in Figure S6.

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The global carbon budget averaged over the last decade (2014–2023) is shown in Figure 2, Figure J6 (right panel) and Table 7. For this period, 90% of the total emissions ( $E_{\text{FOS}} + E_{\text{LUC}}$ ) were from fossil CO<sub>2</sub> emissions ( $E_{\text{FOS}}$ ), and 10% from land-use change ( $E_{\text{LUC}}$ ). The total emissions were partitioned among the atmosphere (48%), ocean (26%) and land (20%), with a small negative budget imbalance ( $\sim 4\%$ ,  $0.4 \text{ GtC yr}^{-1}$ ). For single years, the budget imbalance can be larger (Figure 4). For 2023, the combination of our estimated sources ( $11.1 \pm 0.9 \text{ GtC yr}^{-1}$ ) and sinks ( $11.1 \pm 0.9 \text{ GtC yr}^{-1}$ ) leads to a B<sub>IM</sub> of  $-0.02 \text{ GtC}$ , suggesting a near perfect closure of the global carbon budget.

### 3.9.2 Trend and Variability in the Carbon Budget Imbalance

The carbon budget imbalance (B<sub>IM</sub>; Eq. 1, Figure 4) quantifies the mismatch between the estimated total emissions and the estimated changes in the atmosphere, land, and ocean reservoirs. The budget imbalance from 1960 to 2023 is very small ( $0.5 \text{ GtC}$  over the period, i.e.  $\leq 0.01 \text{ GtC yr}^{-1}$  on average) and shows no trend over the full time series (Figure 4e). The process models (GOBMs and DGVMs) and CO<sub>2</sub>-products have been selected to match observational constraints in the 1990s, but no further constraints have been applied to their representation of trend and variability. Therefore, the near-zero mean and trend in the budget imbalance is seen as evidence of a coherent community understanding of the emissions and their partitioning on those time scales (Figure 4). However, the budget imbalance shows substantial variability of the order of  $\pm 1 \text{ GtC yr}^{-1}$ , particularly over semi-decadal time scales, although most of the variability is within the uncertainty of the estimates. The positive carbon imbalance during the 1960s, and early 1990s, indicates that either the emissions were overestimated, or the sinks were underestimated during these periods. The reverse is true for the 1970s, and to a lesser extent for the 1980s and 2014–2023 period (Figure 4, Table 7).

We cannot attribute the cause of the variability in the budget imbalance with our analysis, we only note that the budget imbalance is unlikely to be explained by errors or biases in the emissions alone because of its large semi-decadal variability component, a variability that is atypical of emissions and has not changed in the past 60 years despite a near tripling in emissions (Figure 4). Errors in S<sub>LAND</sub> and S<sub>OCEAN</sub> are more likely to be the main cause for the budget imbalance, especially on interannual to semi-decadal timescales. For example, underestimation of the S<sub>LAND</sub> by DGVMs has been reported following the eruption of Mount Pinatubo in 1991 possibly due to missing responses to changes in diffuse radiation (Mercado et al., 2009). Although since GCB2021 we accounted for aerosol effects on solar radiation quantity and quality (diffuse vs direct), most DGVMs only used the former as input (i.e., total solar radiation) (Table S1). Thus, the ensemble mean may not capture the full effects of volcanic eruptions, i.e. associated with high light scattering sulphate aerosols, on the land carbon sink (O’Sullivan et al., 2021). DGVMs are suspected to overestimate the land sink in response to the wet decade of the 1970s (Sitch et al., 2008). Quasi-decadal variability in the ocean sink has also been reported, with all methods agreeing on a smaller than expected ocean CO<sub>2</sub> sink in the 1990s and a larger than expected sink in the 2000s (Figure J1; Landschützer et al., 2016, DeVries et al., 2019, Hauck et al., 2020, McKinley et al., 2020, Gruber et al., 2023) and the climate-driven variability could be substantial but is not well constrained (DeVries et al., 2023, Müller et al., 2023). Errors in sink estimates could also be driven by errors in the climatic forcing data, particularly precipitation for S<sub>LAND</sub> and wind for S<sub>OCEAN</sub>. Also, the B<sub>IM</sub> shows substantial departure from

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4117 zero on yearly time scales (Figure 4e), highlighting unresolved variability of the carbon cycle, likely in the land  
4118 sink ( $S_{\text{LAND}}$ ), given its large year to year variability (Figure 4d and 9).

4119 Both the budget imbalance ( $B_{\text{IM}}$ , Table 7) and the residual land sink from the global budget ( $E_{\text{FOS}} + E_{\text{LUC}} - G_{\text{ATM}} -$   
4120  $S_{\text{OCEAN}}$ , Table 5) include an error term due to the inconsistencies that arises from combining  $E_{\text{LUC}}$  from  
4121 bookkeeping models with  $S_{\text{LAND}}$  from DGVMs, most notably the loss of additional sink capacity (see Section  
4122 2.10 and Supplement S.6.4). Other differences include a better accounting of land use changes practices and  
4123 processes in bookkeeping models than in DGVMs, or the bookkeeping models error of having present-day  
4124 observed carbon densities fixed in the past. That the budget imbalance shows no clear trend towards larger  
4125 values over time is an indication that these inconsistencies probably play a minor role compared to other errors  
4126 in  $S_{\text{LAND}}$  or  $S_{\text{OCEAN}}$ .

4127 Although the budget imbalance is near zero for the recent decades, it could be due to a compensation of errors.  
4128 We cannot exclude an overestimation of  $\text{CO}_2$  emissions, particularly from land-use change, given their large  
4129 uncertainty, as has been suggested elsewhere (Piao et al., 2018), and/or an underestimate of the sinks. A larger  
4130 DGVM estimate of the atmosphere-land  $\text{CO}_2$  flux ( $S_{\text{LAND}} - E_{\text{LUC}}$ ) over the extra-tropics would reconcile model  
4131 results with inversion estimates for fluxes in the total land during the past decade (Figure 14; Table 5).

4132 Likewise, a larger  $S_{\text{OCEAN}}$  is also possible given the higher estimates from the  $f\text{CO}_2$ -products, [inversions and](#)  
4133 [oxygen based estimates](#) (see Section 3.6.2, Figure 11 and Figure 14), the underestimation of interior ocean  
4134 anthropogenic carbon accumulation in the GOBMs (Section 3.6.5, [Müller et al., 2023](#)), [known biases of ocean](#)  
4135 [models](#) (e.g., [Terhaar et al., 2022; 2024](#)), [the role of potential temperature bias and skin effects in  \$f\text{CO}\_2\$ -](#)  
4136 [products](#) ([Watson et al., 2020; Dong et al., 2022; Bellenger et al., 2023, Figure 11](#)) and [regionally larger](#)  
4137 [estimates based e.g. on eddy covariance measurements and aircraft data](#) ([Dong et al., 2024a; Long et al., 2021;](#)  
4138 [Jin et al., 2024](#)). More integrated use of observations in the Global Carbon Budget, either on their own or for  
4139 further constraining model results, should help resolve some of the budget imbalance (Peters et al., 2017a).

#### 4140 4 Tracking progress towards mitigation targets

4141 The average growth in global fossil  $\text{CO}_2$  emissions peaked at nearly +3% per year during the 2000s, driven by  
4142 the rapid growth in emissions in China. In the last decade, however, the global growth rate has slowly declined,  
4143 reaching a low +0.6% per year over 2014-2023. While this slowdown in global fossil  $\text{CO}_2$  emissions growth is  
4144 welcome, global fossil  $\text{CO}_2$  emissions continue to grow, far from the rapid emission decreases needed to be  
4145 consistent with the temperature goals of the Paris Agreement.

4146 Since the 1990s, the average growth rate of fossil  $\text{CO}_2$  emissions has continuously declined across the group of  
4147 developed countries of the Organisation for Economic Co-operation and Development (OECD), with emissions  
4148 peaking in around 2005 and declining at 1.4%  $\text{yr}^{-1}$  in the decade 2014-2023, compared to a decline of 0.9%  $\text{yr}^{-1}$   
4149 during the 2004-2013 period (Table 9). In the decade 2014-2023, territorial fossil  $\text{CO}_2$  emissions decreased  
4150 significantly (at the 95% confidence level) in 22 countries/economies whose economies grew significantly (also  
4151 at the 95% confidence level): Belgium, Czechia, Denmark, Estonia, Finland, France, Germany, Jordan,  
4152 Luxembourg, Netherlands, New Zealand, Norway, Portugal, South Korea, Romania, Slovenia, Somalia, Spain,

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Deleted: ), and the recently suggested upward adjustments of the ocean carbon sink in Earth System Models (Terhaar et al., 2022), and in  $f\text{CO}_2$ -products, here related to a potential temperature bias and skin effects (Watson et al., 2020; Dong et al., 2022; Figure 10). If  $S_{\text{OCEAN}}$  were to be based on  $f\text{CO}_2$ -products alone, with all  $f\text{CO}_2$ -products including this adjustment, this would result in a 2013-2022  $S_{\text{OCEAN}}$  of 3.7  $\text{GtC yr}^{-1}$  (Dong et al., 2022) or >3.9  $\text{GtC yr}^{-1}$  (Watson et al., 2020), i.e., outside of the range supported by the atmospheric inversions and with an implied negative  $B_{\text{IM}}$  of more than -1  $\text{GtC yr}^{-1}$  indicating that a closure of the budget could only be achieved with either anthropogenic emissions being significantly larger and/or the net land sink being substantially smaller than estimated here. A recent model study suggests that the skin effect is smaller (about 0.1  $\text{GtC yr}^{-1}$  or 5%) due to feedbacks with surface carbon concentration (Bellenger et al., 2023), which would nevertheless lead to a larger  $S_{\text{OCEAN}}$  even in the GOBMs.

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Sweden, Switzerland, United Kingdom, USA (updated from Le Quéré et al., 2019). Altogether, these 22 countries emitted 2.2 GtC yr<sup>-1</sup> (8.1 GtCO<sub>2</sub> yr<sup>-1</sup>) on average over the last decade, about 23% of world CO<sub>2</sub> fossil emissions. For comparison, 18 countries showed a significant decrease in territorial fossil CO<sub>2</sub> emissions over the previous decade (2004-2013).

Decomposing emission changes into the components of growth, a Kaya decomposition, helps give an initial understanding of the drivers of the changes (Peters et al., 2017b). The reduction in growth in global fossil CO<sub>2</sub> emissions in the last decade is due to slightly weaker economic growth, accelerating declines in CO<sub>2</sub> emissions per unit energy, and sustained declines in energy per unit GDP (Figure 17). These trends are a supposition of the trends at the national level. Fossil CO<sub>2</sub> emission declines in the USA and the EU27 are primarily driven by slightly weaker economic growth since the Global Financial Crisis (GFC) in 2008/2009, sustained declines in energy per GDP, and sustained declines in CO<sub>2</sub> emissions per unit energy with a slight acceleration in the USA in the last decade. In contrast, fossil CO<sub>2</sub> emissions continue to grow in non-OECD countries, although the growth rate has slowed from 4.9% yr<sup>-1</sup> during the 2004-2013 decade to 1.8% yr<sup>-1</sup> in the last decade (Table 9).

Representing 47% of non-OECD emissions in 2023, a large part of this slowdown is due to China, which has seen emissions growth decline from 7.5% yr<sup>-1</sup> in the 2004-2013 decade to 1.9% yr<sup>-1</sup> in the last decade. Excluding China, non-OECD emissions grew at 3% yr<sup>-1</sup> in the 2004-2013 decade, compared to 1.7% yr<sup>-1</sup> in the last decade. China has had weaker economic growth in the 2000s compared to the 2010s, and the rate of reduction in the energy intensity of economic production has weakened significantly since 2015 with accelerating declines in CO<sub>2</sub> emissions per unit energy (Figure 17). India has had strong economic growth that is not offset by declines in energy per GDP or declines in CO<sub>2</sub> emissions per unit energy, driving up fossil CO<sub>2</sub> emissions. Despite the high deployment of renewables in some countries (e.g., China, India), fossil energy sources continue to grow to meet growing energy demand (Le Quéré et al., 2019). In the rest of the world, economic growth has slowed considerably in the last decade, but is only partly offset by declines in energy or carbon intensity, leading to growing emissions.

Globally, fossil CO<sub>2</sub> emissions growth is slowing, and this is due in part to the emergence of climate policy (Eskander and Fankhauser 2020; Le Quere et al 2019) and technological change, which is leading to a shift from coal to gas and growth in renewable energies, and reduced expansion of coal capacity. At the aggregated global level, decarbonisation shows a strong and growing signal in the last decade, with smaller contributions from lower economic growth and declines in energy per GDP (Figure 17). Altogether, global fossil CO<sub>2</sub> emissions are still growing (average of 0.6% per year over the 2014-2023 decade), far from the reductions needed to meet the ambitious climate goals of the UNFCCC Paris agreement.

Last, we update the remaining carbon budget (RCB) based on two studies, the IPCC AR6 (Canadell et al., 2021) and the revision of the IPCC AR6 estimates (Forster et al., 2024, Lamboll et al., 2023). We update the RCB assessed by the IPCC AR6 (Canadell et al., 2021), accounting for the 2020 to 2024 estimated emissions from fossil fuel combustion (E<sub>FOS</sub>) and land use changes (E<sub>LUC</sub>). From January 2025, the IPCC AR6 RCB (50% likelihood) for limiting global warming to 1.5°C, 1.7°C and 2°C is estimated to amount to 85, 180, and 315 GtC (305, 655, 1155 GtCO<sub>2</sub>). The Forster et al. (2024) study proposed a significantly lower RCB than IPCC AR6,

**Deleted:** , Zimbabwe... (updated from Le Quéré et al., 2019). Altogether, these 26... countries emitted 2.7... GtC yr<sup>-1</sup> (10.0... GtCO<sub>2</sub> yr<sup>-1</sup>) on average over the last decade, about 28...3% of world CO<sub>2</sub> fossil emissions. For comparison, 22...8 countries showed a significant decrease in territorial fossil CO<sub>2</sub> emissions over the previous decade (2003-2012). Figure 16 shows that the emission declines in the USA and the EU27 are primarily driven by slightly weaker economic growth in the last decade compared to the 1990s, sustained declines in energy per GDP (though, weakening in the USA), and sustained declines in CO<sub>2</sub> emissions per unit energy (decarbonisation) with a slight acceleration in the USA in the last decade. (... [81])

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with the largest reduction being due to an update of the climate emulator (MAGICC) used to estimate the warming contribution of non-CO<sub>2</sub> agents, and to the warming (i.e. emissions) that occurred over the 2020-2023 period. We update the Forster et al., budget accounting for the 2024 estimated emissions from fossil fuel combustion (EFOS) and land use changes (ELUC). From January 2025, the Forster et al., (2024) RCB (50% likelihood) for limiting global warming to 1.5°C, 1.7°C and 2°C is estimated to amount to 45, 140, and 290 GtC (160, 510, 1060 GtCO<sub>2</sub>), significantly smaller than the updated IPCC AR6 estimate. Both the original IPCC AR6 and Forster et al. (2024) estimates include the Earth System uncertainty on the climate response to cumulative CO<sub>2</sub> emissions, which is reflected through the percent likelihood of exceeding the given temperature threshold, an additional uncertainty of 220GtCO<sub>2</sub> due to alternative non-CO<sub>2</sub> emission scenarios, and other sources of uncertainties (see Canadell et al., 2021). The two sets of estimates overlap when considering all uncertainties. The IPCC AR6 estimates have the advantage of a consensus building approach, while the Forster et al. (2024) estimates include significant update estimates but without the backing of the IPCC yet.

Here, we take the average of our 2024 update of both IPCC AR6 and Forster et al. (2024) estimates, giving a remaining carbon (50% likelihood) for limiting global warming to 1.5°C, 1.7°C and 2°C of respectively 65, 160, and 305 GtC (235, 585, 1110 GtCO<sub>2</sub>) starting from January 2025. We emphasise the large uncertainties, particularly when close to the global warming limit of 1.5°C. These 1.5°C, 1.7°C and 2°C remaining carbon budgets correspond respectively to about 6, 14 and 27 years from the beginning of 2025, at the 2024 level of total anthropogenic CO<sub>2</sub> emissions. Reaching net-zero CO<sub>2</sub> emissions by 2050 entails cutting total anthropogenic CO<sub>2</sub> emissions by about 0.4 GtC (1.6 GtCO<sub>2</sub>, 3.9% of 2024 emissions, each year on average, comparable to the decrease in EFOS observed in 2020 during the COVID-19 pandemic. However, this would lead to cumulative emissions over 2025-2050 of 145 GtC (530 GtCO<sub>2</sub>), well above the remaining carbon budget of 65 GtC to limit global warming to 1.5°C, but still within the remaining budget of 160 GtC to limit warming to 1.7°C (in phase with the “well below 2°C” ambition of the Paris Agreement). Even reaching net zero CO<sub>2</sub> globally by 2040, which would require annual emissions cuts of 0.7 GtC (2.5 GtCO<sub>2</sub>) on average, would still exceed the remaining carbon budget, with 90 GtC (325 GtCO<sub>2</sub>) cumulative emissions over 2025-2040, unless the global emissions trajectory becomes net negative (i.e. more anthropogenic CO<sub>2</sub> sinks than emissions) after 2040.

## 5 Discussion

Each year when the global carbon budget is published, each flux component is updated for all previous years to consider corrections that are the result of further scrutiny and verification of the underlying data in the primary input datasets. Annual estimates may be updated with improvements in data quality and timeliness (e.g., to eliminate the need for extrapolation of forcing data such as land-use). Of all terms in the global budget, only the fossil CO<sub>2</sub> emissions and the growth rate in atmospheric CO<sub>2</sub> concentration are based primarily on empirical inputs supporting annual estimates in this carbon budget. The carbon budget imbalance, yet an imperfect measure, provides a strong indication of the limitations in observations, in understanding and representing processes in models, and/or in the integration of the carbon budget components.

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The persistent unexplained variability in the carbon budget imbalance limits our ability to verify reported emissions (Peters et al., 2017a) and suggests we do not yet have a complete understanding of the underlying carbon cycle dynamics on annual to decadal timescales. Resolving most of this unexplained variability should be possible through different and complementary approaches. First, as intended with our annual updates, the imbalance as an error term should be reduced by improvements of individual components of the global carbon budget that follow from improving the underlying data and statistics and by improving the models through the resolution of some of the key uncertainties detailed in Table 10. Second, additional clues to the origin and processes responsible for the variability in the budget imbalance could be obtained through a closer scrutiny of carbon variability in light of other Earth system data (e.g., heat balance, water balance), and the use of a wider range of biogeochemical observations to better understand the land-ocean partitioning of the carbon imbalance such as the constraint from atmospheric oxygen included this year. Finally, additional information could also be obtained through better inclusion of process knowledge at the regional level, and through the introduction of inferred fluxes such as those based on satellite xCO<sub>2</sub> retrievals. The limit of the resolution of the carbon budget imbalance is yet unclear, but most certainly not yet reached given the possibilities for improvements that lie ahead.

Estimates of global fossil CO<sub>2</sub> emissions from different datasets are in relatively good agreement when the different system boundaries of these datasets are considered (Andrew, 2020a). But while estimates of E<sub>FOS</sub> are derived from reported activity data requiring much fewer complex transformations than some other components of the budget, uncertainties remain, and one reason for the apparently low variation between datasets is precisely the reliance on the same underlying reported energy data. The budget excludes some sources of fossil CO<sub>2</sub> emissions, which available evidence suggests are relatively small (<1%). We have added emissions from lime production in China and the US, but these are still absent in most other non-Annex I countries, and before 1990 in other Annex I countries.

Estimates of E<sub>LUC</sub> suffer from a range of intertwined issues, including the poor quality of historical land-cover and land-use change maps, the rudimentary representation of management processes in most models, and the confusion in methodologies and boundary conditions used across methods (e.g., Arneth et al., 2017; Pongratz et al., 2014, see also Supplement S.6.4 on the loss of sink capacity; Bastos et al., 2021). Uncertainties in current and historical carbon stocks in soils and vegetation also add uncertainty in the E<sub>LUC</sub> estimates. Unless a major effort to resolve these issues is made, little progress is expected in the resolution of E<sub>LUC</sub>. This is particularly concerning given the growing importance of E<sub>LUC</sub> for climate mitigation strategies, and the large issues in the quantification of the cumulative emissions over the historical period that arise from large uncertainties in E<sub>LUC</sub>.

By adding the DGVMs estimates of CO<sub>2</sub> fluxes due to environmental change from countries' managed forest areas (part of S<sub>LAND</sub> in this budget) to the budget E<sub>LUC</sub> estimate, we successfully reconciled the large gap between our E<sub>LUC</sub> estimate and the land use flux from NGHGs using the approach described in Grassi et al. (2021) for future scenarios and in Grassi et al. (2023) using data from the Global Carbon Budget 2021. The updated data presented here can be used as potential adjustment in the policy context, e.g., to help assess the collective countries' progress towards the goal of the Paris Agreement and avoiding double-accounting for the

sink in managed forests. In the absence of this adjustment, collective progress would hence appear better than it is (Grassi et al., 2021). The application of this adjustment is also recommended in the UNFCCC Synthesis report for the first Global Stocktake (UNFCCC, 2022) whenever a comparison between LULUCF fluxes reported by countries and the global emission estimates of the IPCC is conducted. However, this adjustment should be seen as a short-term and pragmatic fix based on existing data, rather than a definitive solution to bridge the differences between global models and national inventories. Additional steps are needed to understand and reconcile the remaining differences, some of which are relevant at the country level (Grassi, et al., 2023, Schwingshackl, et al., 2022).

The comparison of GOBMs,  $f\text{CO}_2$ -products, and inversions highlights substantial discrepancy in the temporal evolution of  $\text{SOCEAN}$  in the Southern Ocean and northern high-latitudes (Figure 14, Hauck et al., 2023a) and in the mean  $\text{SOCEAN}$  in the tropics. A large part of the uncertainty in the mean fluxes stems from the regional distribution of the river flux adjustment term. The current distribution simulates the largest share of the outgassing to occur in the tropics (Lacroix et al., 2020). The long-standing sparse data coverage of  $f\text{CO}_2$  observations in the Southern compared to the Northern Hemisphere (e.g., Takahashi et al., 2009) continues to exist (Bakker et al., 2016, 2024, Figure S1) and to lead to substantially higher uncertainty in the  $\text{SOCEAN}$  estimate for the Southern Hemisphere (Watson et al., 2020, Gloege et al., 2021, Hauck et al., 2023a). This discrepancy, which also hampers model improvement, points to the need for increased high-quality  $f\text{CO}_2$  observations especially in the Southern Ocean. At the same time, model uncertainty is illustrated by the large spread of individual GOBM estimates (indicated by shading in Figure 14) and highlights the need for model improvement. The issue of diverging trends in  $\text{SOCEAN}$  from different methods is smaller this year as the trend in the  $f\text{CO}_2$ -products was revised downwards with the data available in this GCB release, but remains a matter of concern. Recent and on-going work suggests that the  $f\text{CO}_2$ -products may overestimate the trend (Hauck et al., 2023a, Supplement section S3.4), though the full  $f\text{CO}_2$ -product ensemble remains to be tested. A data-constrained model approach suggests that the GOBMs underestimate the amplitude of decadal variability, but that the  $f\text{CO}_2$ -products overestimate the trend (Mayot et al., 2024). At the same time, evidence is accumulating that GOBMs likely underestimate the mean flux (Section 3.6.2, Terhaar et al., 2022, DeVries et al., 2023, Müller et al., 2023, Dong et al., 2024). The independent constraint from atmospheric oxygen measurements gives a larger sink for the past decade and a steeper trend. However, the estimate is consistent within uncertainties with  $\text{SOCEAN}$ , with the relatively larger ocean sink in the  $f\text{CO}_2$ -products and some of the GOBMs. The assessment of the net land-atmosphere exchange from DGVMs and atmospheric inversions also shows substantial discrepancy, particularly for the estimate of the net land flux over the northern extra-tropic. This discrepancy highlights the difficulty to quantify complex processes ( $\text{CO}_2$  fertilisation, nitrogen deposition and fertilisers, climate change and variability, land management, etc.) that collectively determine the net land  $\text{CO}_2$  flux. Resolving the differences in the Northern Hemisphere land sink will require the consideration and inclusion of larger volumes of observations.

We provide metrics for the evaluation of the ocean and land models and the atmospheric inversions (Figures S2 to S4, Table S11). These metrics expand the use of observations in the global carbon budget, helping 1) to support improvements in the ocean and land carbon models that produce the sink estimates, and 2) to constrain

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the representation of key underlying processes in the models and to allocate the regional partitioning of the CO<sub>2</sub> fluxes. The introduction of process-based metrics targeted to evaluate the simulation of SOCEAN in the ocean biogeochemistry models is an important addition to the evaluation based on ocean carbon observations. This is an initial step towards the introduction of a broader range of observations and more stringent model evaluation that we hope will support continued improvements in the annual estimates of the global carbon budget.

We assessed before that a sustained decrease of –1% in global emissions could be detected at the 66% likelihood level after a decade only (Peters et al., 2017a). Similarly, a change in behaviour of the land and/or ocean carbon sink would take as long to detect, and much longer if it emerges more slowly. To continue reducing the carbon imbalance on annual to decadal time scales, regionalising the carbon budget, and integrating multiple variables are powerful ways to shorten the detection limit and ensure the research community can rapidly identify issues of concern in the evolution of the global carbon cycle under the current rapid and unprecedented changing environmental conditions.

## 6 Conclusions

The estimation of global CO<sub>2</sub> emissions and sinks is a major effort by the carbon cycle research community that requires a careful compilation and synthesis of measurements, statistical estimates, and model results. The delivery of an annual carbon budget serves two purposes. First, there is a large demand for up-to-date information on the state of the anthropogenic perturbation of the climate system and its underpinning causes. A broad stakeholder community relies on the datasets associated with the annual carbon budget including scientists, policy makers, businesses, journalists, and non-governmental organisations engaged in adapting to and mitigating human-driven climate change. Second, over the last decades we have seen unprecedented changes in the human and biophysical environments (e.g., changes in the growth of fossil fuel emissions, impact of COVID-19 pandemic, Earth’s warming, and strength of the carbon sinks), which call for frequent assessments of the state of the planet, a better quantification of the causes of changes in the contemporary global carbon cycle, and an improved capacity to anticipate its evolution in the future. Building this scientific understanding to meet the extraordinary climate mitigation challenge requires frequent, robust, transparent, and traceable datasets and methods that can be scrutinised and replicated. This paper via ‘living data’ helps to keep track of new budget updates.

## 7 Data availability

The data presented here are made available in the belief that their wide dissemination will lead to greater understanding and new scientific insights of how the carbon cycle works, how humans are altering it, and how we can mitigate the resulting human-driven climate change. Full contact details and information on how to cite the data shown here are given at the top of each page in the accompanying database and summarised in Table 2.

The accompanying database includes three Excel files organised in the following spreadsheets:

File Global\_Carbon\_Budget\_2024v1.0.xlsx includes the following:

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4525	2. The global carbon budget (1959-2023);	Deleted: 2022
4526	3. The historical global carbon budget (1750-2023);	Deleted: 2022
4527	4. Global CO <sub>2</sub> emissions from fossil fuels and cement production by fuel type, and the per-capita emissions	
4528	(1850-2023);	Deleted: 2022
4529	5. CO <sub>2</sub> emissions from land-use change from the individual bookkeeping models (1959-2023);	Deleted: 2022
4530	6. Ocean CO <sub>2</sub> sink from the individual global ocean biogeochemistry models and fCO <sub>2</sub> -products (1959-	
4531	2023);	Deleted: 2022
4532	7. Terrestrial CO <sub>2</sub> sink from the individual DGVMs (1959-2023);	Deleted: 2022
4533	8. Cement carbonation CO <sub>2</sub> sink (1959-2023).	Deleted: 2022
4534	File National_Fossil_Carbon_Emissions_2024v1.0.xlsx includes the following:	Deleted: 2023v1
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4536	2. Territorial country CO <sub>2</sub> emissions from fossil fuels and cement production (1850-2023);	Deleted: 2022
4537	3. Consumption country CO <sub>2</sub> emissions from fossil fuels and cement production and emissions transfer from	
4538	the international trade of goods and services (1990-2020) using CDIAC/UNFCCC data as reference;	
4539	4. Emissions transfers (Consumption minus territorial emissions; 1990-2020);	
4540	5. Country definitions.	
4541	File National_LandUseChange_Carbon_Emissions_2024v1.0.xlsx includes the following:	Deleted: 2023v1
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4543	2. Territorial country CO <sub>2</sub> emissions from Land Use Change (1850-2023) from three bookkeeping models;	Deleted: 2022
4544	All three spreadsheets are published by the Integrated Carbon Observation System (ICOS) Carbon Portal and	
4545	are available at <a href="https://doi.org/10.18160/GCP-2024">https://doi.org/10.18160/GCP-2024</a> (Friedlingstein et al., 2024). National emissions data are also	Deleted: 2023
4546	available on Zenodo (Andrew and Peters, 2024), from the Global Carbon Atlas	Deleted: 2023
4547	( <a href="http://www.globalcarbonatlas.org/">http://www.globalcarbonatlas.org/</a> , last access: 28 October 2024) and from Our World in Data	Deleted: 2022
4548	( <a href="https://ourworldindata.org/co2-emissions">https://ourworldindata.org/co2-emissions</a> , last access: 28 October 2024).	Deleted: 9 November 2023
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**Author contributions**

PF, MO, MWJ, RMA, JH, PL, CLQ, HL, ITL, AO, GPP, WP, JP, CS, and SSi designed the study, conducted the analysis, and wrote the paper with input from JGC, PCi and RBJ. RMA, GPP and JIK produced the fossil CO<sub>2</sub> emissions and their uncertainties and analysed the emissions data. MH and GMa provided fossil fuel emission data. JP, TGA, ZQ, and CS provided the bookkeeping land-use change emissions with synthesis by JP and CS. SSm provided the estimates of non-vegetation CDR fluxes. JB, MC, ÖG, NG, TI, TJ, LR, JS, RS, and HTs provided an update of the global ocean biogeochemical models. LMD, ARF, DJF, MG, LG, YI, AJ, CR, AR, JZ, and PC provided an update of the ocean fCO<sub>2</sub>-data products, with synthesis on both streams by JH, PL and NMa. SRA, NRB, MB, CFB, HCB, KC, KE, WE, RAF, TGK, SKL, NL, NMe, NMM, SN, LO, TO, DP, AJS, ST, BT, CN, and RW provided ocean fCO<sub>2</sub> measurements for the year 2023, with synthesis by AO and TS. AA, VA, PCa, THC, JD, CDR, AF, JHe, AKJ, EK, JK, PCM, LM, TN, MO, QS, HTi, XYa, WY, XYu, and SZ provided an update of the Dynamic Global Vegetation Models, with synthesis by SSi and MO. HL, RSA, OT, and ET provided estimates of land and ocean sinks from Earth System Models, as well as a projection of the atmospheric growth rate for 2024. NC, FC, ARJ, FJ, ZJ, JL, SM, YN, PIP, CR, DY, and NZ provided an updated atmospheric inversion, WP, FC, and ITL developed the protocol and produced the synthesis and evaluation of the atmospheric inversions. RMA provided projections of the 2024 fossil emissions and atmospheric CO<sub>2</sub> growth rate. PL provided the predictions of the 2024 ocean and land sinks. LPC, GCH, KKG, TMR, GRvdW, WX, and ZY provided forcing data for land-use change. FT and GG provided data for the land-use change NGHGI harmonisation. RFK provided key atmospheric CO<sub>2</sub> data. EJM and RFK provided the atmospheric oxygen constraint on surface net carbon sinks. MWJ provided the historical atmospheric CO<sub>2</sub> concentration and growth rate. MO and NB produced the aerosol diffuse radiative forcing for the DGVMs. IH provided the climate forcing data for the DGVMs. PCM provided the evaluation of the DGVMs. MWJ provided the emissions prior for use in the inversion systems. XD provided seasonal emissions data for most recent years for the emission prior. PF, MO and MWJ coordinated the effort, revised all figures, tables, text and numbers to ensure the update was clear from the 2023 edition and in line with the globalcarbonatlas.org.

**Competing interests.**

At least one of the (co-)authors is a member of the editorial board of Earth System Science Data

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## Tables

**Table 1.** Factors used to convert carbon in various units (by convention, Unit 1 = Unit 2 × conversion).

Unit 1	Unit 2	Conversion	Source
GtC (gigatonnes of carbon)	ppm (parts per million) (a)	2.124 (b)	Ballantyne et al. (2012)
GtC (gigatonnes of carbon)	PgC (petagrams of carbon)	1	SI unit conversion
GtCO <sub>2</sub> (gigatonnes of carbon dioxide)	GtC (gigatonnes of carbon)	3.664	44.01/12.011 in mass equivalent
		(a) Measurements of atmospheric CO <sub>2</sub> concentration have units of dry-air mole fraction. 'ppm' is an abbreviation for micromole/mol, dry air.	
(b) The use of a factor of 2.124 assumes that all the atmosphere is well mixed within one year. In reality, only the troposphere is well mixed and the growth rate of CO <sub>2</sub> concentration in the less well-mixed stratosphere is not measured by sites from the NOAA network. Using a factor of 2.124 makes the approximation that the growth rate of CO <sub>2</sub> concentration in the stratosphere equals that of the troposphere on a yearly basis.			

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Ocean and land CO<sub>2</sub> sinks (S<sub>OCEAN</sub> and S<sub>LAND</sub>)

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**Table 3.** Main methodological changes in the global carbon budget since 2020. Methodological changes introduced in one year are kept for the following years unless noted. Empty cells mean there were no methodological changes introduced that year. Table S9 lists methodological changes from the first global carbon budget publication up to 2019.

Publication year	Fossil fuel emissions		LUC emissions	Reservoirs			Other changes
	Global	Country (territorial)		Atmosphere	Ocean	Land	
2020		India's emissions from Andrew (2020: India); Corrections to Netherlands and Aruba and Soviet emissions before 1950 as per Andrew (2020: CO <sub>2</sub> ); China's coal emissions in 2019 derived from official statistics, emissions now shown for EU27 instead of EU28. Projection for 2020 based on assessment of four approaches.	Average of three bookkeeping models; use of 17 DGVMs. Estimate of gross land use sources and sinks provided	Use of six atmospheric inversions	Based on nine models. River flux revised and partitioned NH <sub>2</sub> , Tropics, SH	Based on 17 models	
2021	Projections are no longer an assessment	Official data included for a	ELUC estimate compared to the		Average of means of eight models	Current year prediction of SLAND	

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Friedlingstein et al. (2022a) GCB2021		nt of four approache s.	number of additional countries, new estimates for South Korea, added emissions from lime productio n in China.	estimates adopted in national GHG inventorie s		and means of seven data- products. Current year prediction of SOCEAN using a feed- forward neural network method	using a feed- forward neural network method		
2022				ELUC provided at country level. Revised componen ts decompos ition of ELUC fluxes. Revision of LUC maps for Brazil. New datasets for peat drainage.	Use of nine atmosphe ric inversions	Average of means of ten models and means of seven data- products	Based on 16 models. Revision of LUC maps for Brazil.		
2023				Refined componen ts decompos ition of ELUC. Revision of LUC maps for Indonesia. Use of updated peat drainage estimates.	Use of 14 atmosphe ric inversions : Additional use of 4 Earth System Models to estimate current year CO2	Additional use of 4 Earth System Models and atmosphe ric oxygen method to assess SOCEAN. Regional distributio n of river flux adjustmen t revised.	Based on 20 models. Additional use of 4 Earth System Models and atmosphe ric oxygen method to assess the net atmosphe re-land flux.	Inclusion of an estimate of Carbon Dioxide Removal	
		2024	Inclusion of 2024	Inclusion of 2024	Fourth bookkeepi	Use of 14 atmosphe	Use of 10 GOBMs, 8	Use of 20 DGVMs	

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<a href="#">This study</a>	<a href="#">projections from Carbon Monitor</a>	<a href="#">projections from Carbon Monitor for China, USA, EU27, India, and Rest of the World</a>	<a href="#">net estimate (LUCE). Update in land-use data (HYDE3.4) including revision of LUC maps for China. Updated definition of forest (re-)growth fluxes (consistent with 2nd State of CDR Report).</a>	<a href="#">net inversions models</a>	<a href="#">fCO<sub>2</sub>-products. Added evaluation for fCO<sub>2</sub>-products.</a>		
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**Table 4.** References for the process models, bookkeeping models, ocean data products, and atmospheric inversions. All models and products are updated with new data to the end of year 2023.

Model/data name	Reference	Change from Global Carbon Budget 2023 (Friedlingstein et al., 2023)
<b>Bookkeeping models for land-use change emissions</b>		
BLUE	Hansis et al. (2015)	No change to model, but simulations performed with LUH2-GCB2024 forcing. Update in added peat drainage emissions.
H&C2023	Houghton and Castanho (2023)	No change to model. Data for years after last modelled year (2020) extrapolated based on anomalies in deforestation fires from GFED. Update in added peat drainage emissions.
OSCAR	Gasser et al. (2020)	No change to model, but land-use forcing changed to LUH2-GCB2024 and FRA2020, extrapolated to 2023. Constraining based on GCB2023 data for SLAND over 1960-2022. Update in added peat drainage emissions.
LUCE	Qin et al. (2024)	New model in GCB2024.
<b>Dynamic global vegetation models</b>		
CABLE-POP	Haverd et al. (2018)	Bug fix applied to land use change calculations enabling variable crop and pasture fractions; corrections to the pre-industrial primary forest fraction in Europe; minor parameter changes
CLASSIC	Melton et al. (2020), Asaadi et al. (2018)	Permeable soil depth reduced to 4 m ; 15 soil layers in the top 4 m permeable soil and 5 bed rock layers from 4 m to 62 m; saturated hydraulic conductivity decreases with depth in the permeable soil layers; transpiration occurs from a partially-wet canopy leaves. These changes yield better runoff seasonality and a more realistic partitioning of precipitation into runoff and evapotranspiration.
CLM6.0	Lawrence et al. (2019)	Updates to surface datasets; improvement of roughness length calculation; updates to snow optical properties and snow thermal conductivity; improved excess ice; improved simulation of burial of vegetation by snow; urban updates, including transient urban, urban properties, and air conditioning; improvements to biomass heat storage; tillage and residue removal; crop phenology and planting dates; improvement to irrigation methods; PFT parameter update.

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DLEM	Tian et al. (2015), You et al. (2022)	Incorporate mechanistic representations of dynamic crop growth and development processes, such as crop-specific phenological development, carbon allocation, yield formation, and biological N fixation. Agricultural management practices such as N fertiliser use, irrigation, tillage, manure application, dynamic crop rotation, cover cropping, and genetic improvements are also included (You et al. 2022).
EDv3	Moorcroft et al. (2001), Ma et al. (2022)	Minor bug fixes, updated fire submodule.
ELM	Yang et al. (2023), Burrows et al. (2020)	No change.
IBIS	Xia et al. (2024)	Improved algorithm of leaf area index.
iMAPLE	Yue et al. (2024)	The updated version of YIBs model with dynamic coupling between carbon and water cycles.
ISAM	Jain et al. (2013), Meiyappan et al. (2015), Shu et al. (2020)	Vertically resolved soil biogeochemistry (carbon and nitrogen) module, following Shu et al. (2020).
ISBA-CTRIIP	Delire et al. (2020)	No change.
JSBACH	Mauritsen et al. (2019), Reick et al. (2021)	Minor bug fixes in post-processing.
JULES-ES	Wiltshire et al. (2021), Sellar et al. (2019), Burton et al. (2019)	Minor bug fixes. (Using JULES v7.4)
LPJ-GUESS	Smith et al. (2014)	No change.
LPJml	Schaphoff et al., 2018, von Bloh et al., 2018, Lutz et al., 2019 (tillage), Heinke et al., 2023 (livestock grazing)	No change.
LPJwsl	Poulter et al. (2011) (d)	Minor bug fixes, weighting of fire carbon by GFED to simulate annual cycle.
LPX-Bern	Lienert and Joos (2018)	No change.
OCN	Zaehle and Friend (2010), Zaehle et al. (2011)	No change.
ORCHIDEEv3	Krinner et al. (2005), Zaehle and Friend (2010), Vuichard et al. (2019)	No change.
SDGVM	Woodward and Lomas (2004), Walker et al. (2017)	Parameter adjustment for reducing evaporation from vegetation that intercepted precipitation, as well as other adjustments to the calculation of evapotranspiration; bug fix in output of monthly NEP, NBP, soilr, and rh; bug fix on cLeaf monthly output; further development on gross land-use transitions, tracking of carbon from wood & crop harvest, and tracking of primary & secondary vegetation.
VISIT	Ito and Inatomi (2012), Kato et al. (2013)	No change.

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<b>Intermediate complexity land carbon cycle model</b>		
CARDAMOM	Bloom et al. (2016), Smallman et al. (2021)	No change.
<b>Global ocean biogeochemistry models</b>		
NEMO-PlankTOM12	Wright et al. (2021)	Minor bug fixes, change to salinity restoring and restart file. New atmospheric CO2 input for simulations A and C.
NEMO4.2-PISCES (IPSL)	Aumont et al. (2015)	Switch to the new model version NEMO4.2-PISCES. Update following the new protocol (with 1750 preindustrial CO2 for spin-up). New atmospheric CO2 input for simulations A and C.
MICOM-HAMOCC (NorESM1-OCv1.2)	Schwinger et al. (2016)	No change in model set-up. New atmospheric CO2 file for simulations A and C. Corrected diagnostic output for pco2atm; diagnostic output for sfco2 and spco2 provided at the air-sea interface (not with respect to dry air at 1 atm).
MPIOM-HAMOCC6	Lacroix et al. Global Change Biology 2021	No change; only updated atmosphere CO2 input for A and C experiments and run 1948-2023.
NEMO3.6-PISCESv2-gas (CNRM)	Berthet et al. (2019) Séférian et al. (2019)	Updated simulations using 1750 preindustrial conditions instead of 1850. No change in model configuration. New atmospheric CO2 input for simulations A and C.
FESOM2.1-REcoM3	Gürses et al. (2023)	Updated atmospheric CO2 for simulations A and C.
MOM6-COBALT (Princeton)	Liao et al. (2020)	No change.
CESM-ETHZ	Doney et al. (2009)	Compared to the 2023 submission, the spinup was extended to 1422 years before 1750. Also, starting at 1751 the new atmospheric CO2 concentrations provided by GCB have been used for simulations A and C.
MRI-ESM2-3	Tsujing et al. (2024), Sakamoto et al. (2023)	Iron circulation and its limitation on primary production are introduced. Updated atmospheric CO2 for simulations A and C.
ACCESS (CSIRO)	Law et al. (2017)	No change in model set-up (since GCB2023). Updated atmospheric CO2 for simulations A and C.
<b>fCO2-products</b>		
VLIZ-SOMFFN (former MPI-SOM-FFN)	Landschützer et al. (2016)	Time period 1982-2023; The estimate now covers the full open ocean and coastal domain as well as the Arctic Ocean extension by merging 2 MLD proxies for year round full coverage. Additionally, in the SOM step, the Seaflux climatology is now used

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Jena-MLS	Rödenbeck et al. (2014) <a href="#">updated to Rödenbeck et al (2022)</a>	<a href="#">Time</a> period extended to <a href="#">2023</a>
<a href="#">CMEMS-LSCE-FFNNv2</a>	<a href="#">Chau et al. (2022)</a>	<a href="#">Time period now 1985-2023</a>
<a href="#">UEXP-FNN-U (previously Watson et al.)</a>	Watson et al. (2020) <a href="#">and Ford et al. (accepted)</a>	<a href="#">Updated CCI-SST to v3 (Embury et al. 2024), with cool bias with respect to global drifter observations corrected following recommendations in Dong et al. (2022). Updated SOM-FFN implementation to Python.</a>
<a href="#">NIES-ML3</a>	<a href="#">Zeng et al. (2022)</a>	<a href="#">Updated time period 1982-2023.</a>
<a href="#">JMA-MLR</a>	<a href="#">Iida et al. (2021)</a>	<a href="#">Time period extended to 2023</a>
<a href="#">OceanSODA-ETHZv2</a>	<a href="#">Gregor et al. (2024)</a>	<a href="#">Updated method following Gregor et al 2024 and time period extended to 2023</a>
<a href="#">LDEO-HPD</a>	<a href="#">Gloege et al. 2022 and Bennington et al. 2022</a>	<a href="#">Timeperiod extended to 2023</a>
<a href="#">CSIR-ML6</a>	<a href="#">Gregor et al. (2019)</a>	<a href="#">Time period 1982-2023.</a>
<b>Atmospheric inversions</b>		
Jena CarboScope	Rödenbeck et al. (2018), <a href="#">Stephens et al. (2007)</a>	<a href="#">Extension to end of year 2023. Slight change in station set. In the NBE-T inversion, removal of the relaxation term, instead, filtering out decadal variations from air temperature. Adding an additive correction to the result of the NBE-T inversion; to account for CO2 flux IAV not related to air temperature, based on 8 long atmospheric records available near-continuously since at least 1976. TM3 driven by ERA5 rather than NCEP.</a>
CAMS	Chevallier et al. (2005), Remaud et al. (2018)	<a href="#">Extension to year 2023. Increase of the 3D resolution with hexagonal prisms rather than rectangular parallelepipeds (3 times more 3D cells than the previous submission). Update of the prior fluxes.</a>
CarbonTracker Europe (CTE)	van der Laan-Luijkx et al. (2017)	<a href="#">Extension to 2023, update of prior fluxes.</a>
NISMON-CO2	Niwa et al. (2022), <a href="#">Niwa et al. (2017)</a>	<a href="#">Extension to 2023, update of prior fluxes.</a>
CT-NOAA	<a href="#">Jacobson et al. (2023a)</a> , <a href="#">Jacobson et al. (2024)</a> , <a href="#">Byrne et al. (2023)</a> , <a href="#">Krol et al. (2005)</a> , <a href="#">Peiro et al. (2022)</a>	<a href="#">Extended to 2023 using the CarbonTracker Near-Real Time release CT-NRT.v2024-1 (Jacobson et al. 2024).</a>
CMS-Flux	Liu et al. (2021)	<a href="#">Extension to 2023, update of prior fluxes.</a>
CAMS-Satellite	Chevallier et al. (2005), Remaud et al. (2018)	<a href="#">Extension to year 2023. Increase of the 3D resolution with hexagonal prisms rather than rectangular parallelepipeds (3 times more 3D cells than the previous submission). Update of the prior fluxes.</a>

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GONGGA	Jin et al. (2023), <a href="#">Nassar et al. (2010)</a>	<a href="#">Extension to 2023, update of prior fluxes.</a>
COLA	Liu et al. (2022)	<a href="#">Extension to 2023, update of prior fluxes.</a>
GCASv2	Jiang et al. (2021) & <a href="#">Emmons et al. (2010)</a>	<a href="#">Extension to 2023, update of prior fluxes.</a>
UoE in-situ	Feng et al. (2016) & Palmer et al. (2019)	<a href="#">Extension to 2023, update of prior fluxes.</a>
IAPCAS	<a href="#">Yang et al. (2021)</a> & <a href="#">Feng et al. (2016)</a>	<a href="#">Extension to 2023, update of prior fluxes.</a>
MIROC4-ACTM	Chandra et al. (2022) & <a href="#">Patra et al. (2018)</a>	<a href="#">Extension to 2023, update of prior fluxes using only CASA and not VISIT. Less observation sites used in the assimilation (46 instead of 50).</a>
<a href="#">NTFVAR</a>	<a href="#">Navagam et al. (2024)</a> & <a href="#">Maksyutov et al. (2021)</a>	<a href="#">New this year</a>
<b>Earth System Models</b>		
CanESM5	Swart et al. (2019), Sospedra-Alfonso et al. (2021)	<a href="#">Reconstructions are extended to 1960-2023, and predictions are extended to 2024.</a>
<a href="#">EC-Earth3-CC</a>	<a href="#">Döscher et al. (2021)</a> , <a href="#">Bilbao et al. (2021)</a> , <a href="#">Bernardello et al. (2024)</a>	<a href="#">New this year.</a>
<a href="#">IPSL-CM6A-CO2-LR</a>	<a href="#">Boucher et al. (2020)</a>	<a href="#">Reconstructions are extended to 1960-2023, and predictions are extended to 2024. No change to model, the CMIP6 CovidMIP CO2 emissions after 2015 are used.</a>
MIROC-ES2L	Watanabe et al. (2020)	<a href="#">Reconstructions are extended to 1960-2023, and predictions are extended to 2024. No change to model, the simulations were rerun including a long spinup.</a>
MPI-ESM1-2-LR	Mauritsen et al. (2019), Li et al. (2023)	<a href="#">Reconstructions are extended to 1960-2023, and predictions are extended to 2024.</a>

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**Table 5.** Comparison of results from the bookkeeping method and budget residuals with results from the DGVMs, as well as additional estimates from atmospheric oxygen, atmospheric inversions and Earth System Models (ESMs) for different periods, the last decade, and the last year available. All values are in GtCyr<sup>-1</sup>. See Figure 7 for explanation of the bookkeeping component fluxes. The DGVM uncertainties represent  $\pm 1\sigma$  of the decadal or annual (for 2023) estimates from the individual DGVMs: for the inverse systems the mean and range of available results is given. All values are rounded to the nearest 0.1 GtC and therefore columns do not necessarily add to zero.

		<u>Mean (GtC/yr)</u>						
		1960s	1970s	1980s	1990s	2000s	2010s	2023
Land-use change emissions (ELUC)	Bookkeeping (BK) Net flux (1a)	1.6±0.7	1.4±0.7	1.4±0.7	1.6±0.7	1.4±0.7	1.1±0.7	1±0.7
	BK - deforestation (total)	1.7 [1.3, 2.1]	1.6 [1.2, 2.0]	1.6 [1.2, 2.0]	1.8 [1.4, 2.2]	1.9 [1.5, 2.3]	1.7 [1.3, 2.1]	1.7 [1.3, 2.1]
	BK - forest regrowth (total)	-0.8 [-1.1, -0.6]	-0.9 [-1.2, -0.6]	-0.9 [-1.2, -0.6]	-1.1 [-1.4, -0.8]	-1.2 [-1.5, -0.9]	-1.2 [-1.5, -0.9]	-1.2 [-1.5, -0.9]
	BK - other transitions	0.3 [0.2, 0.4]	0.2 [0.1, 0.3]	0.2 [0.1, 0.3]	0.4 [0.2, 0.6]	0.4 [0.2, 0.6]	0.1 [0.0, 0.2]	0 [0.0, 0.1]
	BK - peat drainage & peat fires	0.2 [0.1, 0.2]	0.2 [0.1, 0.2]	0.2 [0.1, 0.2]	0.3 [0.2, 0.4]	0.2 [0.1, 0.3]	0.2 [0.1, 0.3]	0.2 [0.2, 0.3]
	BK - wood harvest & forest management	0.2 [-0.2, 0.6]	0.3 [-0.2, 0.7]	0.3 [-0.2, 0.7]	0.3 [-0.2, 0.7]	0.3 [-0.2, 0.7]	0.3 [-0.2, 0.7]	0.3 [0.0, 0.7]
	DGVMs-net flux (1b)	1.5±0.5	1.5±0.5	1.5±0.5	1.7±0.5	1.7±0.5	1.5±0.5	1.2±0.7
Terrestrial sink (SLAND)	Residual sink from global budget (E <sub>FOS</sub> +E <sub>ELUC</sub> (1a)-G <sub>ATM</sub> -S <sub>OCEAN</sub> ) (2a)	1.7±0.8	1.9±0.8	1.6±0.9	2.6±0.9	2.8±0.9	2.7±0.9	2.3±1
	DGVMs (2b)	1.2±0.5	2±0.8	1.8±0.8	2.5±0.6	2.8±0.7	3.2±0.9	2.3±1
Net land fluxes (SLAND-ELUC)	GCB2024 Budget (2b-1a)	-0.4±0.9	0.5±1.1	0.4±1.1	0.9±0.9	1.4±1.1	2.1±1.1	1.3±1.2
	Atmospheric O <sub>2</sub>	---	---	---	1.3±0.7	1±0.7	1±0.8	-
	DGVMs-net (2b-1b)	-0.3±0.5	0.5±0.7	0.3±0.6	0.8±0.4	1.1±0.4	1.7±0.6	1.1±0.8

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**Table 6:** Comparison of results for the ocean sink from the  $f\text{CO}_2$ -products, from global ocean biogeochemistry models (GOBMs), the best estimate for [GCB2024](#) as calculated from  $f\text{CO}_2$ -products and GOBMs that is used in the budget Table 7, as well as additional estimates from atmospheric oxygen, atmospheric inversions and Earth System Models (ESMs) for different periods, the last decade, and the last year available. All values are in  $\text{GtCyr}^{-1}$ . Uncertainties represent  $\pm 1\sigma$  of the estimates from the GOBMs ( $N > 10$ ) and range of ensemble members is given for ensembles with  $N < 10$  ( $f\text{CO}_2$ -products, inversions, ESMs). The uncertainty of the [GCB2024](#) budget estimate is based on expert judgement (Section 2 and Supplementary S1 to S4) and for oxygen it is the standard deviation of a Monte Carlo ensemble (Section 2.8).

<i>Mean (GtC/yr)</i>									
Product	1960s	1970s	1980s	1990s	2000s	2014-2023	2023		
$f\text{CO}_2$ -products	---	---	---	2.3 [1.9, 2.9]	2.5 [2.3, 3.2]	3.1 [2.9, 3.7]	3 [2.4, 3.4]		
GOBMs	1.1 ± 0.2	1.3 ± 0.3	1.8 ± 0.3	2.1 ± 0.3	2.2 ± 0.3	2.6 ± 0.4	2.7 ± 0.4		
GCB2024 Budget	1.2 ± 0.4	1.5 ± 0.4	1.9 ± 0.4	2.1 ± 0.4	2.3 ± 0.4	2.9 ± 0.4	2.9 ± 0.4		
Atmospheric $\text{O}_2$	---	---	---	2 ± 0.5	2.8 ± 0.4	3.4 ± 0.5	---		
Inversions	- [-]	- [-]	1.8 [1.8, 1.9]	2.3 [2.1, 2.5]	2.5 [2.3, 3.1]	3.1 [2.4, 4.1]	3 [1.4, 4.1]		
ESMs	0.7 [0.7, 1.0]	1 [1.0, 1.2]	---	---	1.4 [1.2, 1.7]	1.7 [1.5, 2.1]	1.9 [1.5, 2.2]	2.5 [2.2, 2.8]	2.5 [2.2, 2.8]

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**Table 7:** Decadal mean in the five components of the anthropogenic CO<sub>2</sub> budget for different periods, and last year available. All values are in GtC yr<sup>-1</sup>, and uncertainties are reported as  $\pm 1\sigma$ . Fossil CO<sub>2</sub> emissions include cement carbonation. The table also shows the budget imbalance (B<sub>IM</sub>), which provides a measure of the discrepancies among the nearly independent estimates. A positive imbalance means the emissions are overestimated and/or the sinks are too small. All values are rounded to the nearest 0.1 GtC and therefore columns do not necessarily add to zero.

		<i>Mean (GtC/yr)</i>							
		1960s	1970s	1980s	1990s	2000s	2014-2023	2023	2024 (Projection)
Total emissions (EFOS + ELUC)	Fossil CO <sub>2</sub> emissions (EFOS)	3±0.2	4.7±0.2	5.5±0.3	6.4±0.3	7.8±0.4	9.7±0.5	10.1±0.5	10.2±0.5
	Land-use change emissions (ELUC)	1.6±0.7	1.4±0.7	1.4±0.7	1.6±0.7	1.4±0.7	1.1±0.7	1.1±0.7	1.1±0.7
	Total emissions	4.6±0.7	6.1±0.7	6.9±0.8	7.9±0.8	9.2±0.8	10.8±0.9	11.1±0.9	11.3±0.9
Partitioning	Growth rate in atmospheric CO <sub>2</sub> (GATM)	1.7±0.07	2.8±0.07	3.4±0.02	3.1±0.02	4±0.02	5.2±0.02	5.9±0.2	5.9±0.5
	Ocean sink (SOCEAN)	1.2±0.4	1.5±0.4	1.9±0.4	2.1±0.4	2.3±0.4	2.9±0.4	2.9±0.4	3±0.6
	Terrestrial sink (SLAND)	1.2±0.5	2±0.8	1.8±0.8	2.5±0.6	2.8±0.7	3.2±0.9	2.3±1	3.2±1.5
Budget imbalance	B <sub>IM</sub> =EFOS+ELUC-(GATM+SOCEAN)	0.5	±0.1	±0.2	0.1	0	-0.4	0	-0.7

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	1960s	1970s	1980s	1990s	2000s	2014-2023	2024 (Projection)
AN+SL AND							

\*Fossil emissions excluding the cement carbonation sink amount to  $3\pm0.2$  GtC/yr,  $4.7\pm0.2$  GtC/yr,  $5.5\pm0.3$  GtC/yr,  $6.4\pm0.3$  GtC/yr,  $7.9\pm0.4$  GtC/yr, and  $9.9\pm0.5$  GtC/yr for the decades 1960s to 2010s respectively and to  $10.3\pm0.5$  GtC/yr for 2023, and  $10.4\pm0.5$  GtC/yr for 2024.

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**Table 8.** Cumulative CO<sub>2</sub> for different time periods in gigatonnes of carbon (GtC). Fossil CO<sub>2</sub> emissions include cement carbonation. The budget imbalance (BIM) provides a measure of the discrepancies among the nearly independent estimates. All values are rounded to the nearest 5 GtC and therefore columns do not necessarily add to zero. Uncertainties are reported as follows: EFOS is 5% of cumulative emissions; ELUC prior to 1959 is 1σ spread from the DGVMs, ELUC post-1959 is 0.7\*number of years (where 0.7 GtC/yr is the uncertainty on the annual ELUC flux estimate); GATM uncertainty is held constant at 5 GtC for all time periods; SOCEAN uncertainty is 20% of the cumulative sink (20% relates to the annual uncertainty of 0.4 GtC/yr, which is ~20% of the current ocean sink); and SLAND is the 1σ spread from the DGVMs estimates.

		1750-2023	1850-2014	1850-2023	1960-2023	1850-2024
Emissions	Fossil CO <sub>2</sub> emissions (EFOS)	490±25	400±20	490±25	410±20	500±25
	Land-use change emissions (ELUC)	255±75	215±60	225±65	90±45	225±65
	Total emissions	745±80	615±65	710±70	500±50	725±70
Partitioning	Growth rate in atmos CO <sub>2</sub> (GATM)	305±5	235±5	285±5	220±5	290±5
	Ocean sink (SOCEAN)	195±40	160±30	185±35	130±25	185±35
	Terrestrial sink (SLAND)	245±65	190±55	220±60	150±40	225±60
Budget imbalance	BIM=EFOS+ELUC-(GATM+SOCEAN+SLAND)	0	30	25	0	20
	ND					

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**Table 9.** Average annual growth rate in fossil CO<sub>2</sub> emissions over the most recent decade (2014-2023) and the previous decade (2004-2013). The data for the World include the cement carbonation sink. IAS are emissions from international aviation and shipping. Rest of the World is World minus China, USA, EU27, India and IAS.

	World	China	USA	EU27	India	OECD	Non-OECD	IAS	Rest of the World
2004-2013	2.4%	7.5%	-1.4%	-1.8%	6.4%	-0.9%	4.9%	2.6%	1.9%
2014-2023	0.6%	1.9%	-1.2%	-2.1%	3.6%	-1.4%	1.8%	-1.6%	0.4%

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**Table 10.** Major known sources of uncertainties in each component of the Global Carbon Budget, defined as input data or processes that have a demonstrated effect of at least  $\pm 0.3 \text{ GtC yr}^{-1}$ .

<u>Source of uncertainty</u>	<u>Time scale (years)</u>	<u>Location</u>	<u>Evidence</u>
<u>Fossil CO<sub>2</sub> emissions (EFOS; Section 2.1)</u>			
<u>energy statistics</u>	<u>annual to decadal</u>	<u>global, but mainly China &amp; major developing countries</u>	<u>(Korsbakken et al., 2016, Guan et al., 2012)</u>
<u>carbon content of coal</u>	<u>annual to decadal</u>	<u>global, but mainly China &amp; major developing countries</u>	<u>(Liu et al., 2015)</u>
<u>system boundary</u>	<u>annual to decadal</u>	<u>all countries</u>	<u>(Andrew, 2020a)</u>
<u>Net land-use change flux (ELUC; section 2.2)</u>			
<u>land-cover and land-use change statistics</u>	<u>continuous</u>	<u>global; in particular tropics</u>	<u>(Houghton et al., 2012, Gasser et al., 2020, Ganzenmüller et al., 2022, Yu et al. 2022)</u>
<u>sub-grid-scale transitions</u>	<u>annual to decadal</u>	<u>global</u>	<u>(Wilkenskjeld et al., 2014, Bastos et al., 2021)</u>
<u>vegetation biomass</u>	<u>annual to decadal</u>	<u>global; in particular tropics</u>	<u>(Houghton et al., 2012, Bastos et al., 2021)</u>
<u>forest degradation (fire, selective logging)</u>	<u>annual to decadal</u>	<u>tropics; Amazon</u>	<u>(Aragão et al., 2018, Qin et al., 2021, Lapola et al., 2023)</u>
<u>wood and crop harvest</u>	<u>annual to decadal</u>	<u>global; SE Asia</u>	<u>(Arneth et al., 2017, Erb et al., 2018)</u>
<u>peat burning</u>	<u>multi-decadal trend</u>	<u>global</u>	<u>(van der Werf et al., 2010, 2017)</u>
<u>loss of additional sink capacity</u>	<u>multi-decadal trend</u>	<u>global</u>	<u>(Pongratz et al., 2014, Gasser et al., 2020; Obermeier et al., 2021; Dorgeist et al., 2024)</u>
<u>environmental effects</u>	<u>multi-decadal trend</u>	<u>global</u>	<u>(Gasser et al. 2020, Dorgeist et al., 2024)</u>
<u>Atmospheric growth rate (GATM; section 2.4) no demonstrated uncertainties larger than <math>\pm 0.3 \text{ GtC yr}^{-1}</math>. The uncertainties in GATM have been estimated as <math>\pm 0.2 \text{ GtC yr}^{-1}</math>, although the conversion of the growth rate into a global annual flux assuming instantaneous mixing throughout the atmosphere introduces additional errors that have not yet been quantified.</u>			
<u>Ocean sink (SOCEAN; section 2.5)</u>			

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<a href="#">sparsity in surface fCO2 observations</a>	<a href="#">mean, decadal variability and trend</a>	<a href="#">global, in particular southern hemisphere</a>	<a href="#">(Gloege et al., 2021, Denvil-Sommer et al., 2021, Hauck et al., 2023a; Dong et al., 2024b)</a>
<a href="#">riverine carbon outgassing and its anthropogenic perturbation</a>	<a href="#">annual to decadal</a>	<a href="#">global, in particular partitioning between Tropics and South</a>	<a href="#">(Aumont et al., 2001, Lacroix et al., 2020, Crisp et al., 2022)</a>
<a href="#">Models underestimate interior ocean anthropogenic carbon storage</a>	<a href="#">annual to decadal</a>	<a href="#">global</a>	<a href="#">(Friedlingstein et al., 2022a, this study, DeVries et al., 2023, Müller et al., 2023)</a>
<a href="#">near-surface temperature and salinity gradients</a>	<a href="#">mean on all time-scales</a>	<a href="#">global</a>	<a href="#">(Watson et al., 2020, Dong et al., 2022, Bellenger et al., 2023, Dong et al., 2024a)</a>
<a href="#">Land sink (SLAND; section 2.6)</a>			
<a href="#">strength of CO2 fertilisation</a>	<a href="#">multi-decadal trend</a>	<a href="#">global</a>	<a href="#">(Wenzel et al., 2016; Walker et al., 2021)</a>
<a href="#">response to variability in temperature and rainfall</a>	<a href="#">annual to decadal</a>	<a href="#">global; in particular tropics</a>	<a href="#">(Cox et al., 2013; Jung et al., 2017; Humphrey et al., 2018; 2021)</a>
<a href="#">nutrient limitation and supply</a>	<a href="#">annual to decadal</a>	<a href="#">global</a>	<a href="#">(Zaehle et al., 2014)</a>
<a href="#">carbon allocation and tissue turnover rates</a>	<a href="#">annual to decadal</a>	<a href="#">global</a>	<a href="#">(De Kauwe et al., 2014; O'Sullivan et al., 2022)</a>
<a href="#">tree mortality</a>	<a href="#">annual</a>	<a href="#">global in particular tropics</a>	<a href="#">(Hubau et al., 2021; Brien et al., 2020)</a>
<a href="#">response to diffuse radiation estimation under constant pre-industrial land cover</a>	<a href="#">annual</a>	<a href="#">global</a>	<a href="#">(Mercado et al., 2009; O'Sullivan et al., 2021)</a>
	<a href="#">multi-decadal trend</a>	<a href="#">global</a>	<a href="#">(Gasser et al. 2020, Dorgeist et al., 2024)</a>

## Figures and Captions

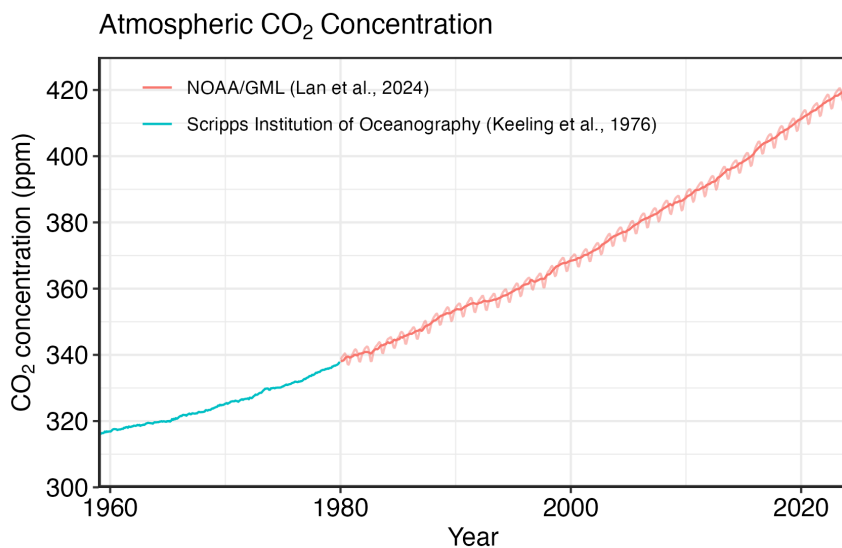
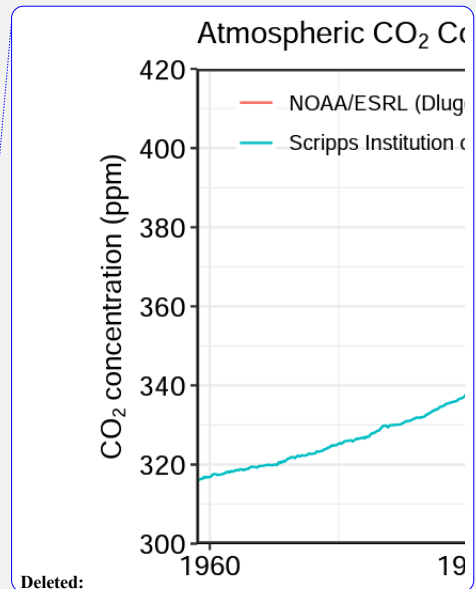


Figure 1. Surface average atmospheric CO<sub>2</sub> concentration (ppm). Since 1980, monthly data are from NOAA/GML (Lan et al., 2024) and are based on an average of direct atmospheric CO<sub>2</sub> measurements from multiple stations in the marine boundary layer (Masarie and Tans, 1995). The 1958-1979 monthly data are from the Scripps Institution of Oceanography, based on an average of direct atmospheric CO<sub>2</sub> measurements from the Mauna Loa and South Pole stations (Keeling et al., 1976). To account for the difference of mean CO<sub>2</sub> and seasonality between the NOAA/GML and the Scripps station networks used here, the Scripps surface average (from two stations) was de-seasonalised and adjusted to match the NOAA/GML surface average (from multiple stations) by adding the mean difference of 0.667 ppm, calculated here from overlapping data during 1980-2012.

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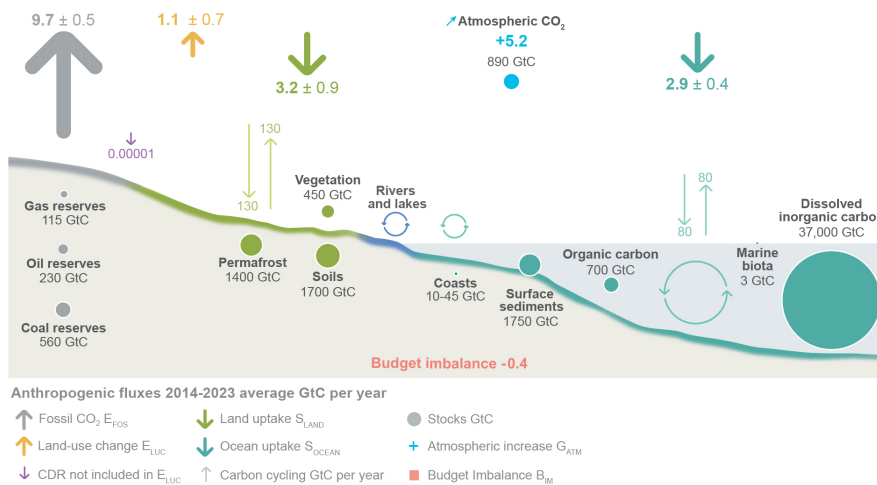


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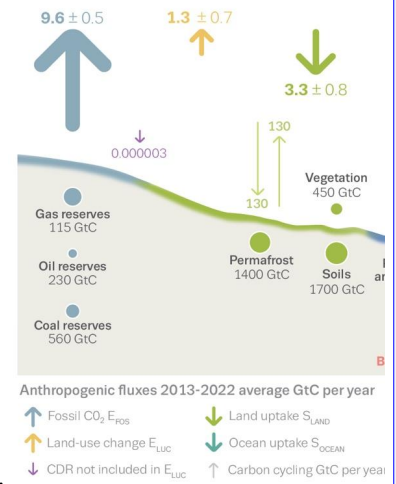
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## The global carbon cycle



**Figure 2.** Schematic representation of the overall perturbation of the global carbon cycle caused by anthropogenic activities, averaged globally for the decade [2014-2023](#). See legends for the corresponding arrows. Fluxes estimates and their 1 standard deviation uncertainty are as reported in Table 7. [The CDR estimate is for the year 2023 only](#). The uncertainty in the atmospheric CO<sub>2</sub> growth rate is very small ( $\pm 0.02$  GtC yr<sup>-1</sup>) and is neglected for the figure. The anthropogenic perturbation occurs on top of an active carbon cycle, with fluxes and stocks represented in the background and taken from Canadell et al. (2021) for all numbers, except for the carbon stocks in coasts which is from a literature review of coastal marine sediments (Price and Warren, 2016). Fluxes are in GtC yr<sup>-1</sup> and reservoirs in GtC.

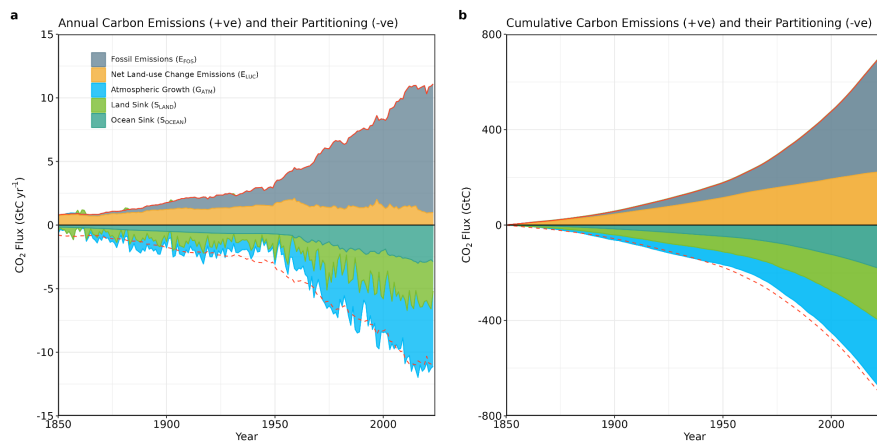
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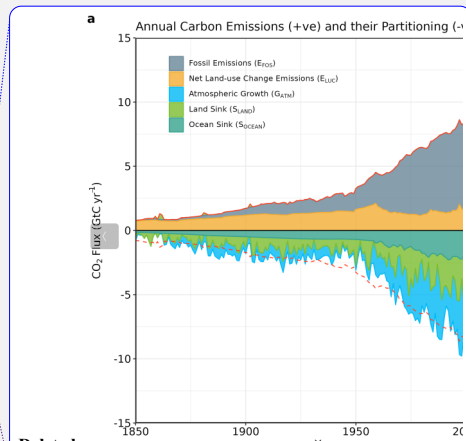
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**Figure 3.** Combined components of the global carbon budget as a function of time, for fossil CO<sub>2</sub> emissions ( $E_{FOS}$ , including a small sink from cement carbonation; grey) and emissions from land-use change ( $E_{LUC}$ ; brown), as well as their partitioning among the atmosphere ( $G_{ATM}$ ; cyan), ocean ( $S_{OCEAN}$ ; blue), and land ( $S_{LAND}$ ; green). Panel (a) shows annual estimates of each flux (in  $GtC\ yr^{-1}$ ) and panel (b) the cumulative flux (the sum of all prior annual fluxes, in  $GtC$ ) since the year 1850. The partitioning is based on nearly independent estimates from observations (for  $G_{ATM}$ ) and from process model ensembles constrained by data (for  $S_{OCEAN}$  and  $S_{LAND}$ ) and does not exactly add up to the sum of the emissions, resulting in a budget imbalance ( $B_{IM}$ ) which is represented by the difference between the bottom red line (mirroring total emissions) and the sum of carbon fluxes in the ocean, land, and atmosphere reservoirs. All data are in  $GtC\ yr^{-1}$  (panel a) and  $GtC$  (panel b). The  $E_{FOS}$  estimate is based on a mosaic of different datasets, and has an uncertainty of  $\pm 5\%$  ( $\pm 1\sigma$ ). The  $E_{LUC}$  estimate is from [four](#) bookkeeping models (Table 4) with uncertainty of  $\pm 0.7\ GtC\ yr^{-1}$ . The  $G_{ATM}$  estimates prior to 1959 are from Joos and Spahni (2008) with uncertainties equivalent to about  $\pm 0.1$ - $0.15\ GtC\ yr^{-1}$  and from Lan et al. (2024) since 1959 with uncertainties of about  $\pm 0.07\ GtC\ yr^{-1}$  during 1959-1979 and  $\pm 0.02\ GtC\ yr^{-1}$  since 1980. The  $S_{OCEAN}$  estimate is the average from Khatiwala et al. (2013) and DeVries (2014) with uncertainty of about  $\pm 30\%$  prior to 1959, and the average of an ensemble of models and an ensemble of  $fCO_2$ -products (Table 4) with uncertainties of about  $\pm 0.4\ GtC\ yr^{-1}$  since 1959. The  $S_{LAND}$  estimate is the average of an ensemble of models (Table 4) with uncertainties of about  $\pm 1\ GtC\ yr^{-1}$ . See the text for more details of each component and their uncertainties.

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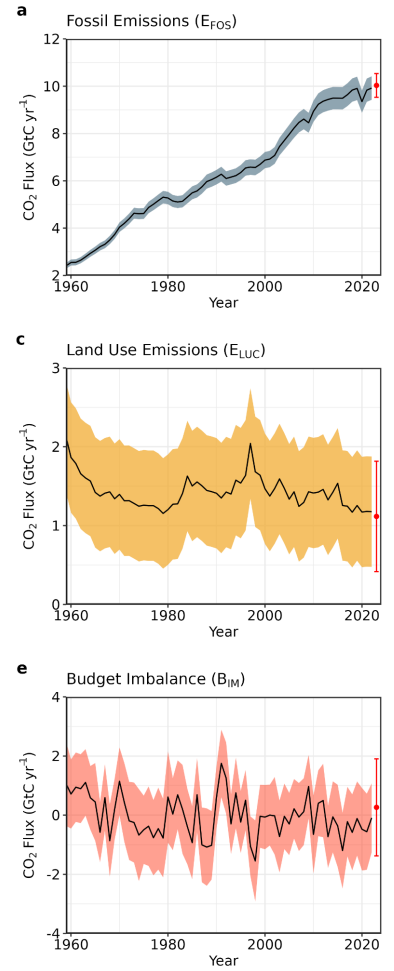
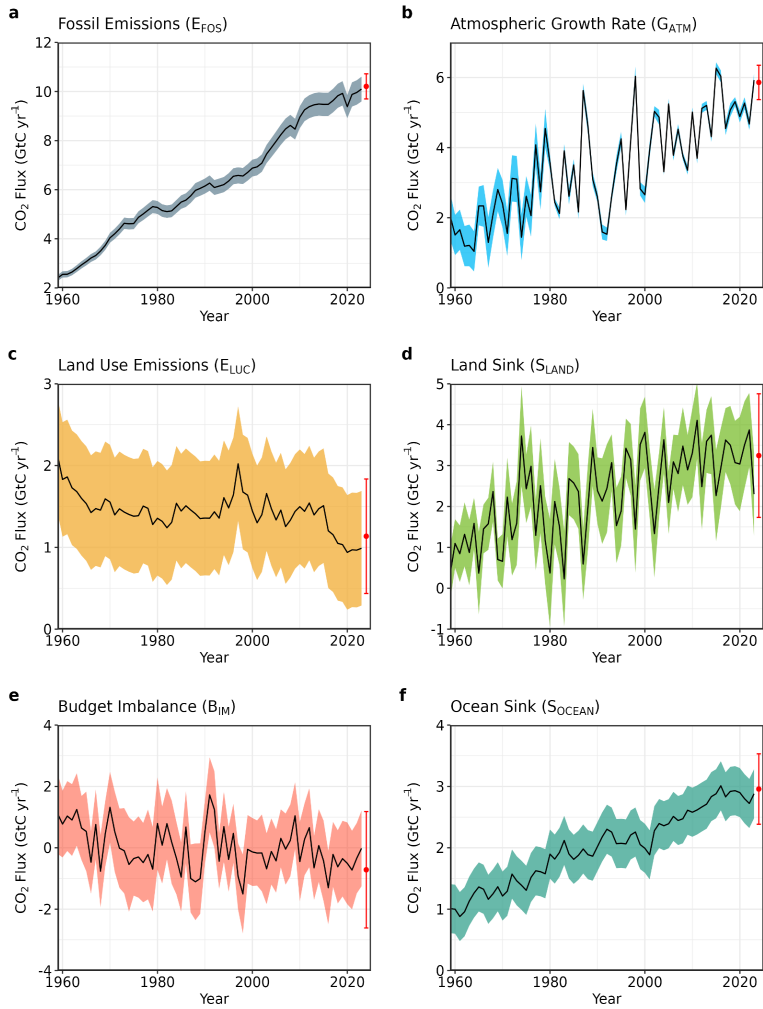
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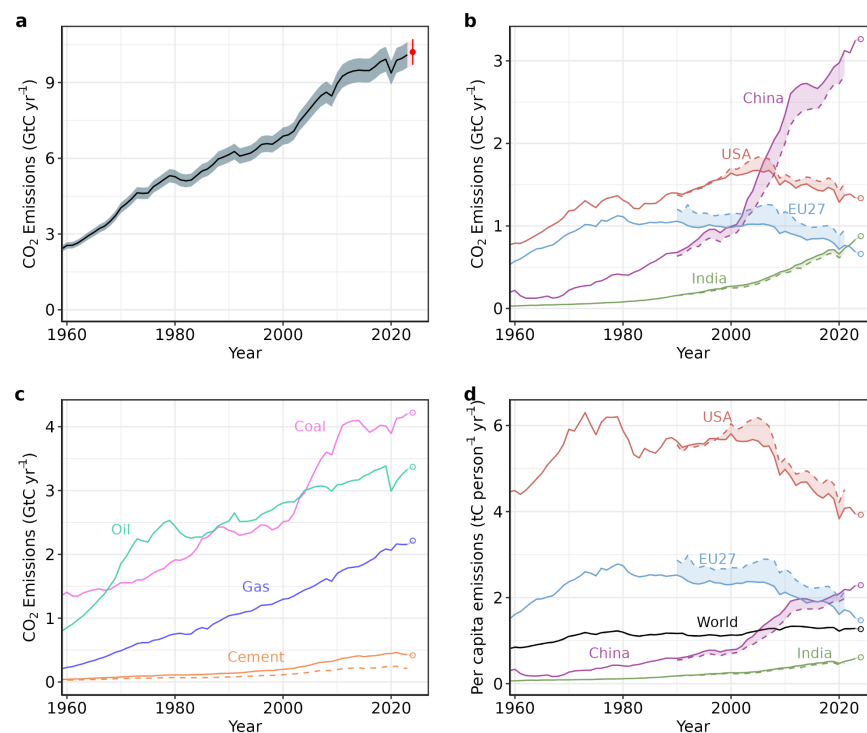
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**Figure 4.** Components of the global carbon budget and their uncertainties as a function of time, presented individually for (a) fossil CO<sub>2</sub> including cement carbonation emissions ( $E_{\text{FOS}}$ ), (b) growth rate in atmospheric CO<sub>2</sub> concentration ( $G_{\text{ATM}}$ ), (c) emissions from land-use change ( $E_{\text{LUC}}$ ), (d) the land CO<sub>2</sub> sink ( $S_{\text{LAND}}$ ), (e) the ocean CO<sub>2</sub> sink ( $S_{\text{OCEAN}}$ ), (f) the budget imbalance ( $B_{\text{IM}}$ ) that is not accounted for by the other terms. Positive values of  $S_{\text{LAND}}$  and  $S_{\text{OCEAN}}$  represent a flux from the atmosphere to land or the ocean. All data are in GtC yr<sup>-1</sup> with the uncertainty bounds representing  $\pm 1$  standard deviation in shaded colour. Data sources are as in Figure 3. The red dots indicate our projections for the year 2024 and the red error bars the uncertainty in the 2024 projections (see methods).

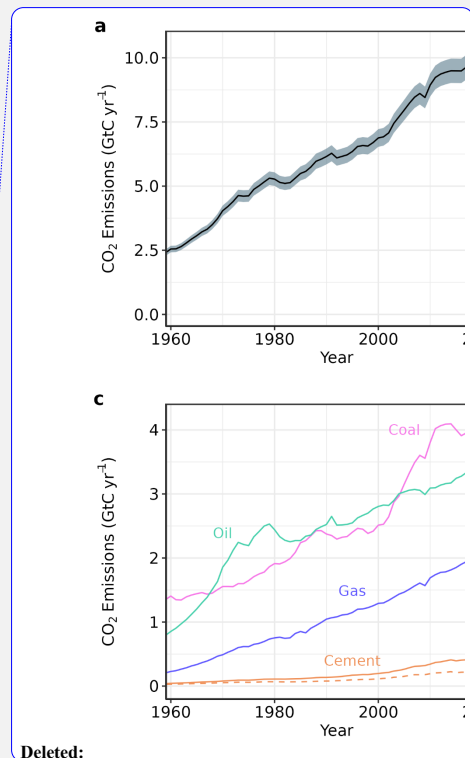


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**Figure 5.** Fossil CO<sub>2</sub> emissions for (a) the globe, including an uncertainty of  $\pm 5\%$  (grey shading) and a projection through the year 2024 (red dot and uncertainty range), (b) territorial (solid lines) and consumption (dashed lines) emissions for the top three country emitters (USA, China, India) and for the European Union (EU27), (c) global emissions by fuel type, including coal, oil, gas, and cement, and cement minus cement carbonation (dashed), and (d) per-capita emissions the world and for the large emitters as in panel (b). Territorial emissions are primarily from a draft update of Hefner and Marland (2023) except for national data for most Annex I countries for 1990–2022, which are reported to the UNFCCC as detailed in the text, as well as some improvements in individual countries, and extrapolated forward to 2023 using data from Energy Institute. Consumption-based emissions are updated from Peters et al. (2011a). See Section 2.1 and Supplement S.1 for details of the calculations and data sources.

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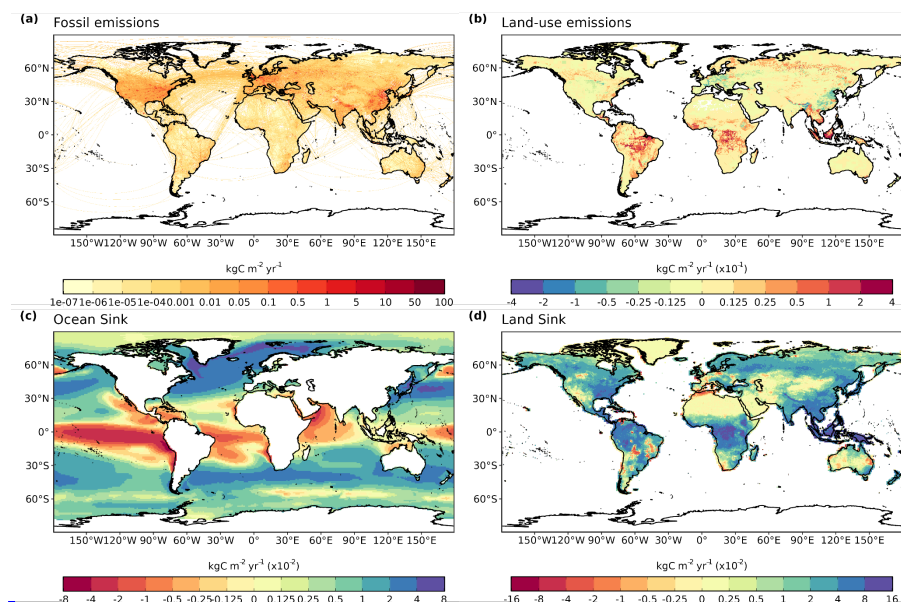
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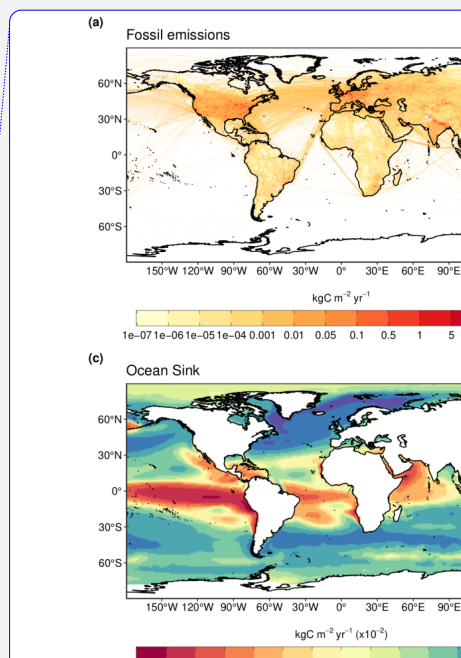
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**Figure 6.** The [2014-2023](#) decadal mean components of the global carbon budget, presented for (a) fossil CO<sub>2</sub> emissions ( $E_{FOS}$ ), (b) land-use change emissions ( $E_{LUC}$ ), (c) the ocean CO<sub>2</sub> sink ( $S_{OCEAN}$ ), and (d) the land CO<sub>2</sub> sink ( $S_{LAND}$ ). Positive values for  $E_{FOS}$  and  $E_{LUC}$  represent a flux to the atmosphere, whereas positive values of  $S_{OCEAN}$  and  $S_{LAND}$  represent a flux from the atmosphere to the ocean or the land (carbon sink). In all panels, yellow/red colours represent a source (flux from the land/ocean to the atmosphere), green/blue colours represent a sink (flux from the atmosphere into the land/ocean). All units are in  $\text{kgC m}^{-2} \text{yr}^{-1}$ . Note the different scales in each panel.  $E_{FOS}$  data shown is from GCP-[GridFEDv2024.0](#) and does not include cement carbonation. The  $E_{LUC}$  map shows the average  $E_{LUC}$  from the [four](#) bookkeeping models plus emissions from peat drainage and peat fires. [BLUE](#) and [LUCE](#) provide spatially explicit estimates at  $0.25^\circ$  resolution. Gridded  $E_{LUC}$  estimates for H&C2023 and OSCAR are derived by spatially distributing their national data based on the spatial patterns of BLUE gross fluxes in each country (see Schwingshackl et al., 2022, for more details about the methodology).  $S_{OCEAN}$  data shown is the average of GOBMs and [CO<sub>2</sub>-products](#) means, using GOBMs simulation A, no adjustment for bias and drift applied to the gridded fields (see Section 2.5).  $S_{LAND}$  data shown is the average of the DGVMs for simulation S2 (see Section 2.6).

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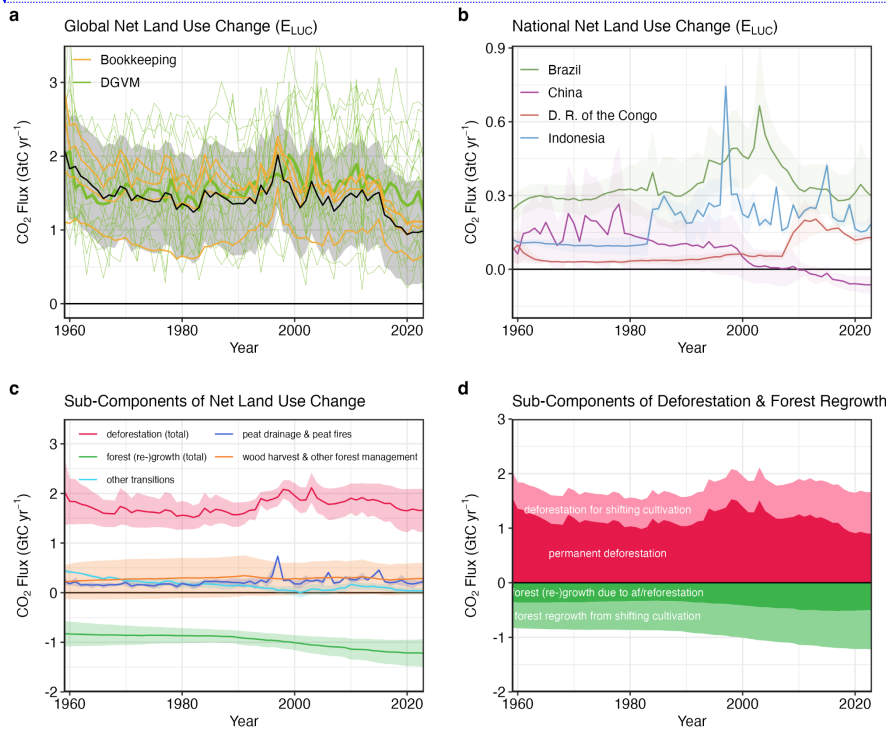
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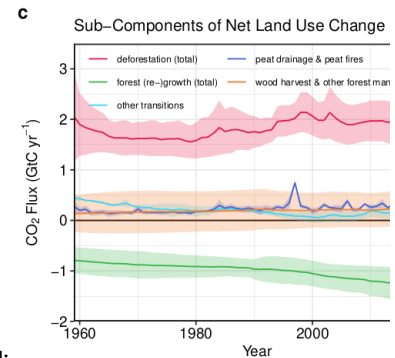
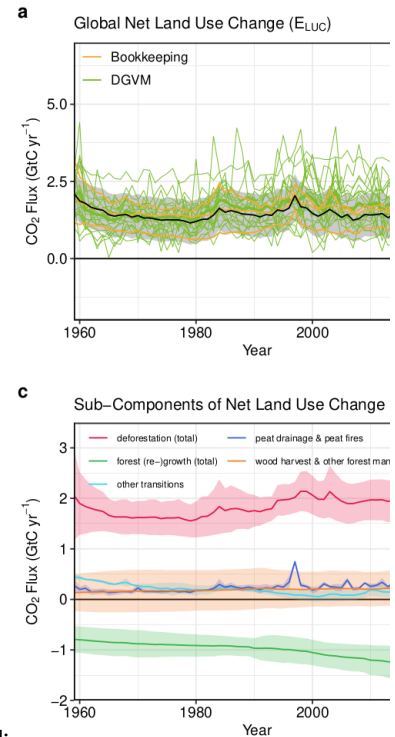
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**Figure 7.** Net CO<sub>2</sub> exchanges between the atmosphere and the terrestrial biosphere related to land use change. (a) Net CO<sub>2</sub> emissions from land-use change ( $E_{LUC}$ ) with estimates from the [four](#) bookkeeping models (yellow lines) and the budget estimate (black with  $\pm 1\sigma$  uncertainty), which is the average of the [four](#) bookkeeping models. Estimates from individual DGVMs (narrow green lines) and the DGVM ensemble mean (thick green line) are also shown. (b) Net CO<sub>2</sub> emissions from land-use change from the four countries with largest cumulative emissions since 1959. Values shown are the average of the [four](#) bookkeeping models, with shaded regions as  $\pm 1\sigma$  uncertainty. (c) Sub-components of  $E_{LUC}$ : (i) emissions from deforestation (including permanent deforestation and deforestation in shifting cultivation cycles), (ii) emissions from peat drainage & peat fires, (iii) removals from forest (re-)growth (including forest (re-)growth due to afforestation and reforestation and forest regrowth in shifting cultivation cycles), (iv) fluxes from wood harvest and other forest management (comprising slash and product decay following wood harvest, regrowth after wood harvest, and fire suppression), and (v) emissions and removals related to other land-use transitions. The sum of the five components is  $E_{LUC}$  shown in panel (a). (d) Sub-components of ‘deforestation (total)’ and of ‘forest (re-)growth (total)’: (i) deforestation in shifting cultivation cycles, (ii) permanent deforestation, (iii) forest (re-)growth due to afforestation and/or reforestation, and (iv) forest regrowth in shifting cultivation cycles.



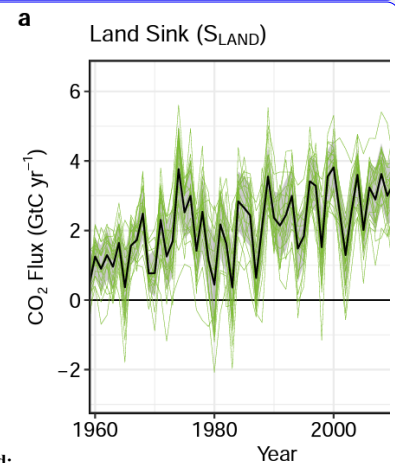
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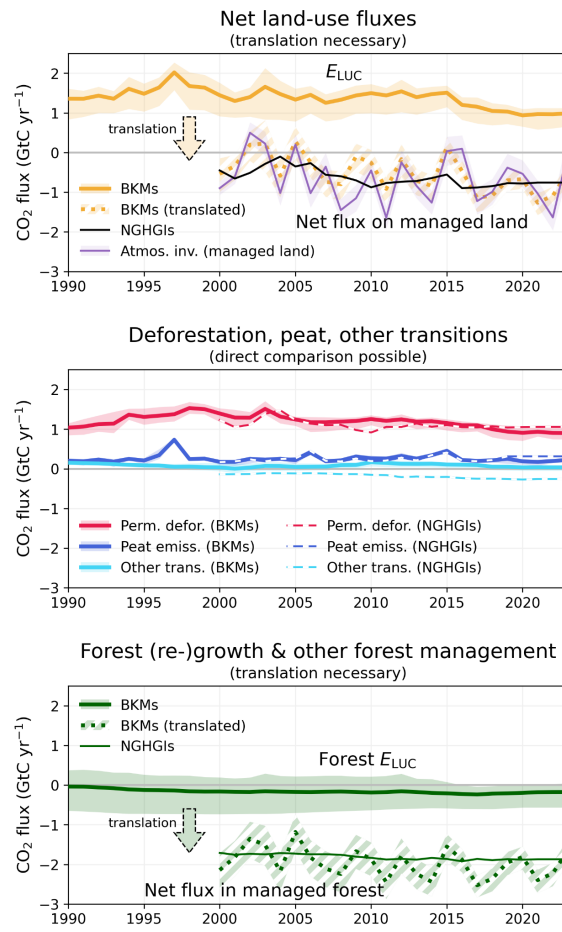
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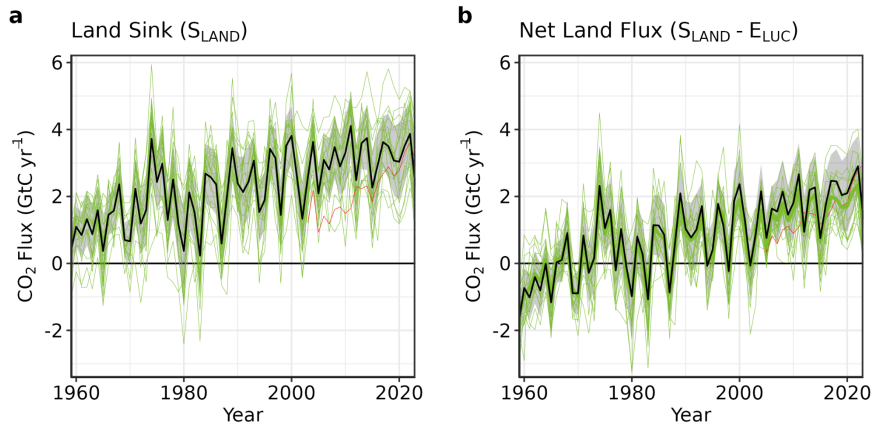
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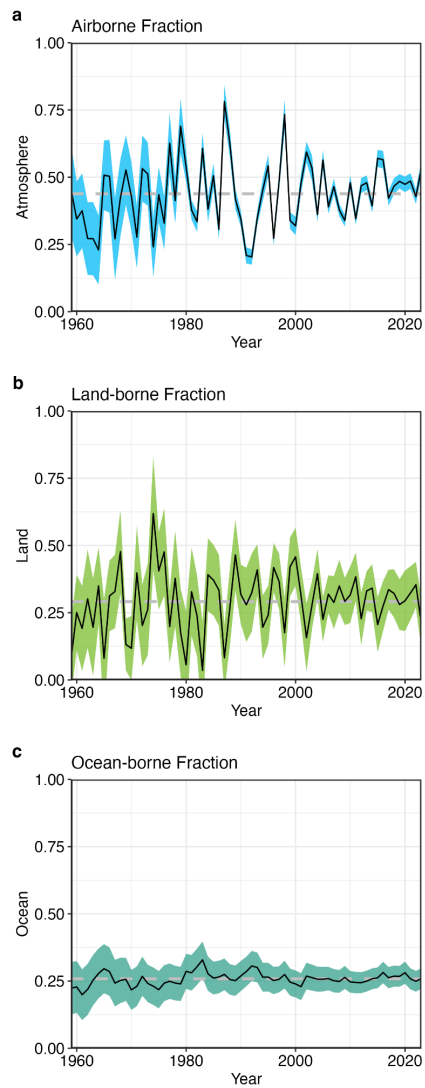
**Figure 8.** Comparison of land-use flux estimates from bookkeeping models (BKMs; following the GCB definition of  $E_{\text{LUC}}$ ), national GHG inventories (NGHGs; following IPCC guidelines and thus including all carbon fluxes on managed land), and atmospheric inversion systems (considering fluxes on managed land only). To compare BKM results with NGHGs, a translation is necessary for some subcomponents. (a) Net land-use fluxes, for which a translation of BKMs is necessary, (b) subcomponents permanent deforestation, peat drainage & peat fires, and other transitions, which can be directly compared and (c) subcomponent forest (re-)growth & other forest management, for which a translation is necessary. The lines represent the mean of 4 BKMs and 14 atmospheric inversion estimates, respectively; Shaded areas denote the full range across BKM estimates and the standard deviation for atmospheric inversions, respectively. The subcomponent forest (re-)growth & other forest management includes removals from forest (re-)growth (permanent), emissions and removals from wood harvest & other forest management, and emissions and removals in shifting cultivation cycles. The translation of

BKM estimates to NGHGI estimates in (a) and (c) is done by adding the natural land sink in managed forests to the BKM estimates (see also Table S10). The GCB definition of ELUC and the NGHGI definition of land-use fluxes are equally valid, each in its own context. For illustrative purposes we only show the translation of BKM estimates to the NGHGI definition.

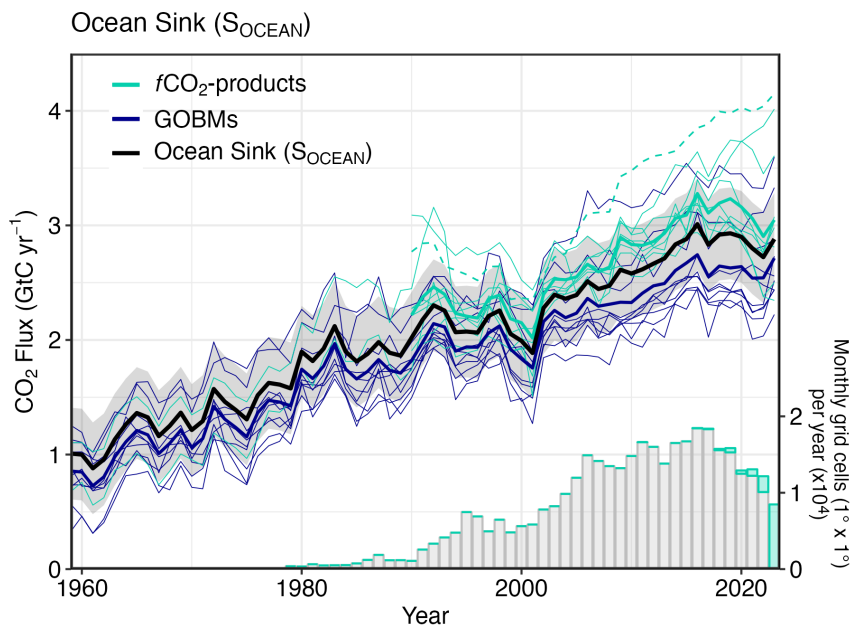
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**Figure 9.** (a) The land CO<sub>2</sub> sink ( $S_{\text{LAND}}$ ) estimated by individual DGVMs (green), and CARDAMOM (red), as well as the budget estimate (black with  $\pm 1\sigma$  uncertainty), which is the average of all DGVMs. (b) Net atmosphere-land CO<sub>2</sub> fluxes ( $S_{\text{LAND}} - E_{\text{LUC}}$ ). The budget estimate of the net land flux (black with  $\pm 1\sigma$  uncertainty) combines the DGVM estimate of  $S_{\text{LAND}}$  from panel (a) with the bookkeeping estimate of  $E_{\text{LUC}}$  from Figure 7a. Uncertainties are similarly propagated in quadrature. DGVMs also provide estimates of  $E_{\text{LUC}}$  (see Figure 7a), which can be combined with their own estimates of the land sink. Hence panel (b) also includes an estimate for the net land flux for individual DGVMs (thin green lines) and their multi-model mean (thick green line).

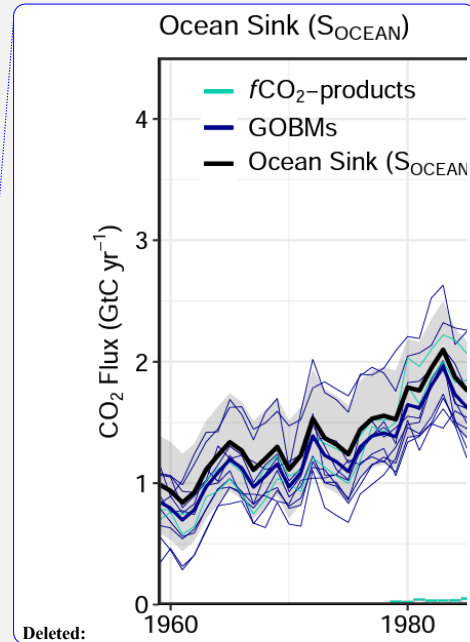


**Figure 10.** The partitioning of total anthropogenic CO<sub>2</sub> emissions (EFOS + ELUC) across (a) the atmosphere (airborne fraction), (b) land (land-borne fraction), and (c) ocean (ocean-borne fraction). Black lines represent the central estimate, and the coloured shading represents the uncertainty. The grey dashed lines represent the long-term average of the airborne (44%), land-borne (30%) and ocean-borne (25%) fractions during 1960–2023 (with a B<sub>IM</sub> of 1%).



**Figure 11.** Comparison of the anthropogenic atmosphere-ocean CO<sub>2</sub> flux showing the budget values of  $S_{\text{OCEAN}}$  (black; with the uncertainty in grey shading), individual ocean models (royal blue), and the ocean  $f\text{CO}_2$ -products (cyan; with UExP-FFN-U, previously Watson et al. (2020), in dashed line as not used for ensemble mean). Two  $f\text{CO}_2$ -products (Jena-MLS, LDEO-HPD) extend back to 1959. The  $f\text{CO}_2$ -products were adjusted for the pre-industrial ocean source of CO<sub>2</sub> from river input to the ocean, by subtracting a source of 0.65 GtC yr<sup>-1</sup> to make them comparable to  $S_{\text{OCEAN}}$  (see Section 2.5). Bar-plot in the lower right illustrates the number of monthly gridded values in the SOCAT v2024 database (Bakker et al., 2024). Grey bars indicate the number of grid cells in SOCAT v2023, and coloured bars indicate the newly added grid cells in v2024.

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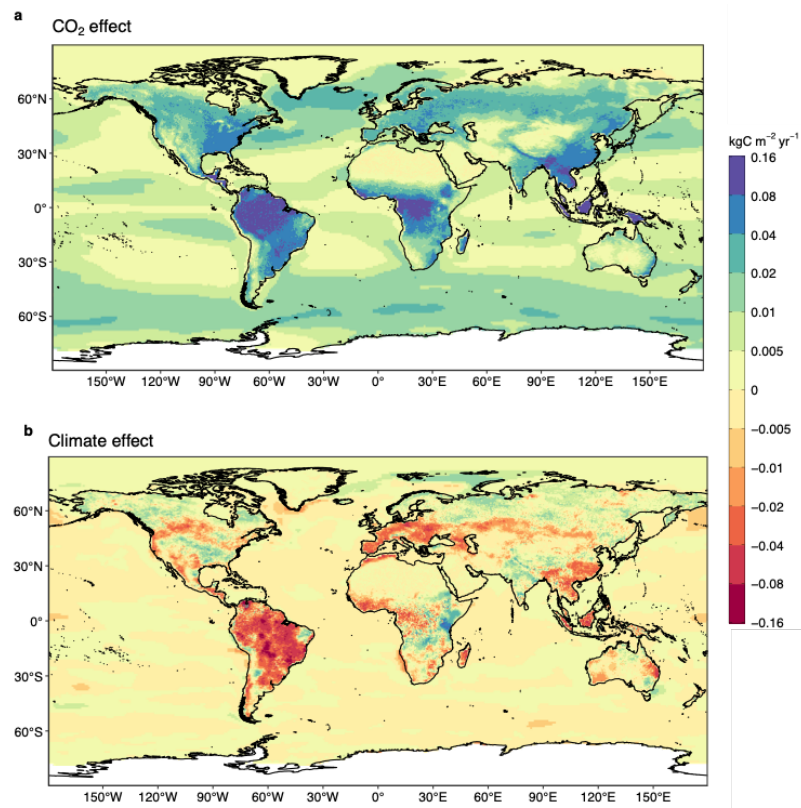
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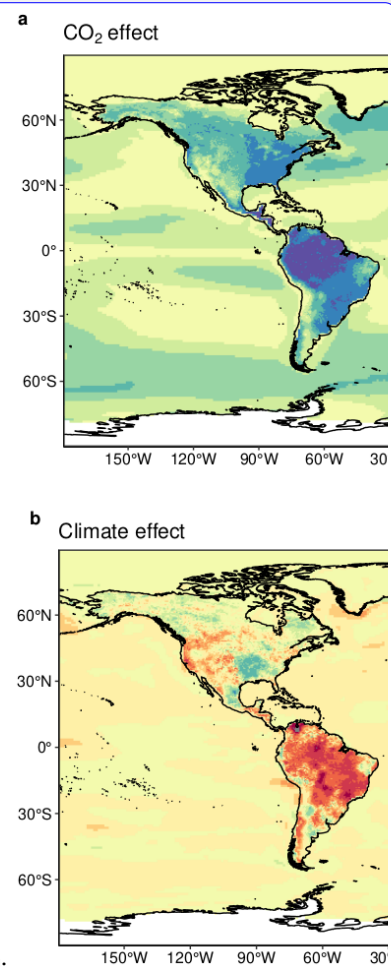
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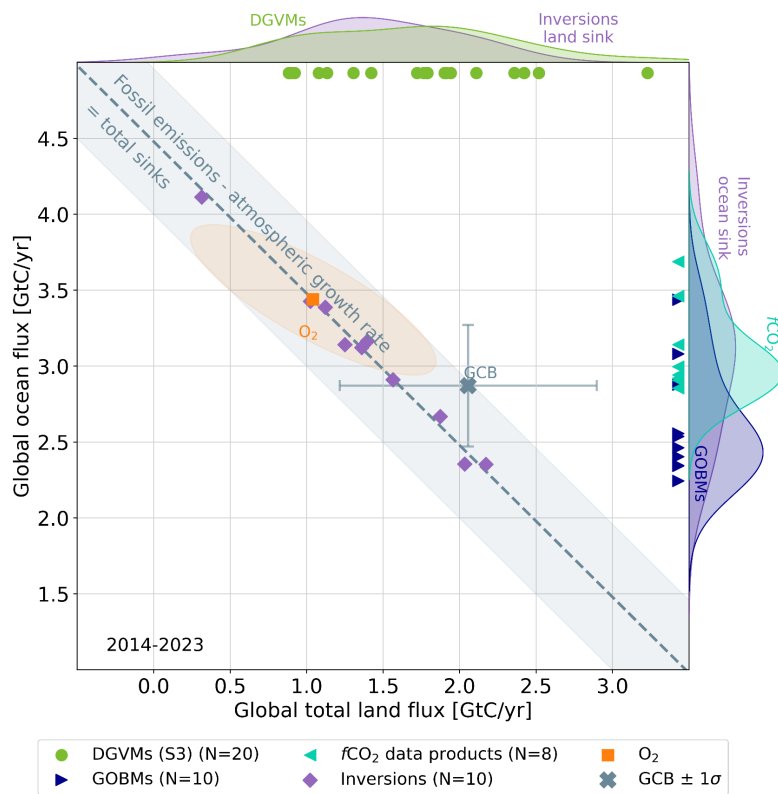
**Figure 12.** Attribution of the atmosphere-ocean ( $S_{OCEAN}$ ) and atmosphere-land ( $S_{LAND}$ ) CO<sub>2</sub> fluxes to (a) increasing atmospheric CO<sub>2</sub> concentrations and (b) changes in climate, averaged over the previous decade [2014-2023](#). All data shown is from the processed-based GOBMs and DGVMS. Note that the sum of ocean CO<sub>2</sub> and climate effects shown here will not equal the ocean sink shown in Figure 6, which includes the  $f_{CO_2}$ -products. See Supplement S.3.2 and S.4.1 for attribution methodology. Units are in kgC m<sup>-2</sup> yr<sup>-1</sup> (note the non-linear colour scale). Positive values (blue) are CO<sub>2</sub> sinks, negative values (red) are CO<sub>2</sub> sources.



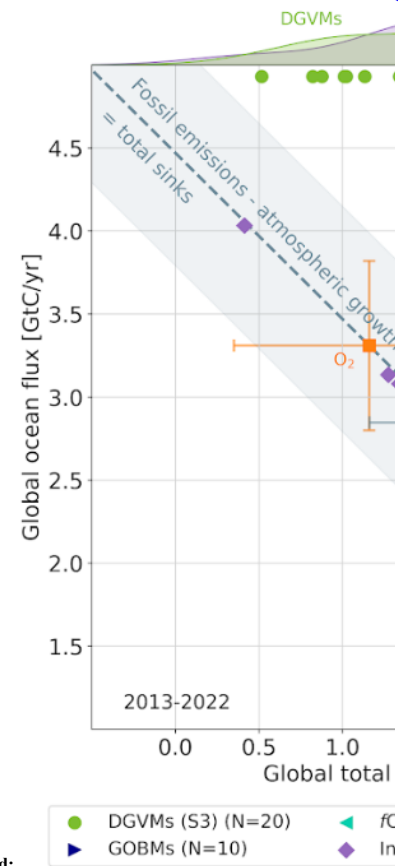
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**Figure 13.** The 2014-2023 decadal mean global net atmosphere-ocean and atmosphere-land fluxes derived from the ocean models and  $f\text{CO}_2$  products (y-axis, right and left pointing blue triangles respectively), and from the DGVMs (x-axis, green symbols), and the same fluxes estimated from the atmospheric inversions (purple symbols). The shaded distributions show the densities of the ensembles of individual estimates. The grey central cross is the mean ( $\pm 1\sigma$ ) of  $S_{\text{OCEAN}}$  and  $(S_{\text{LAND}} - E_{\text{LUC}})$  as assessed in this budget. The grey diagonal line represents the constraint on the global land + ocean net flux, i.e. global fossil fuel emissions minus the atmospheric growth rate from this budget ( $E_{\text{FOS}} - G_{\text{ATM}}$ ). The orange square represents the same global net atmosphere-ocean and atmosphere-land fluxes as estimated from the atmospheric  $\text{O}_2$  constraint (the ellipse drawn around the central atmospheric  $\text{O}_2$  estimate is a contour representing the  $1\sigma$  uncertainty of the land and ocean fluxes as a joint probability distribution). Positive values are  $\text{CO}_2$  sinks. Note that the inverse estimates have been scaled for a minor difference between  $E_{\text{FOS}}$  and GridFEDv2024.0 (Jones et al., 2024a).



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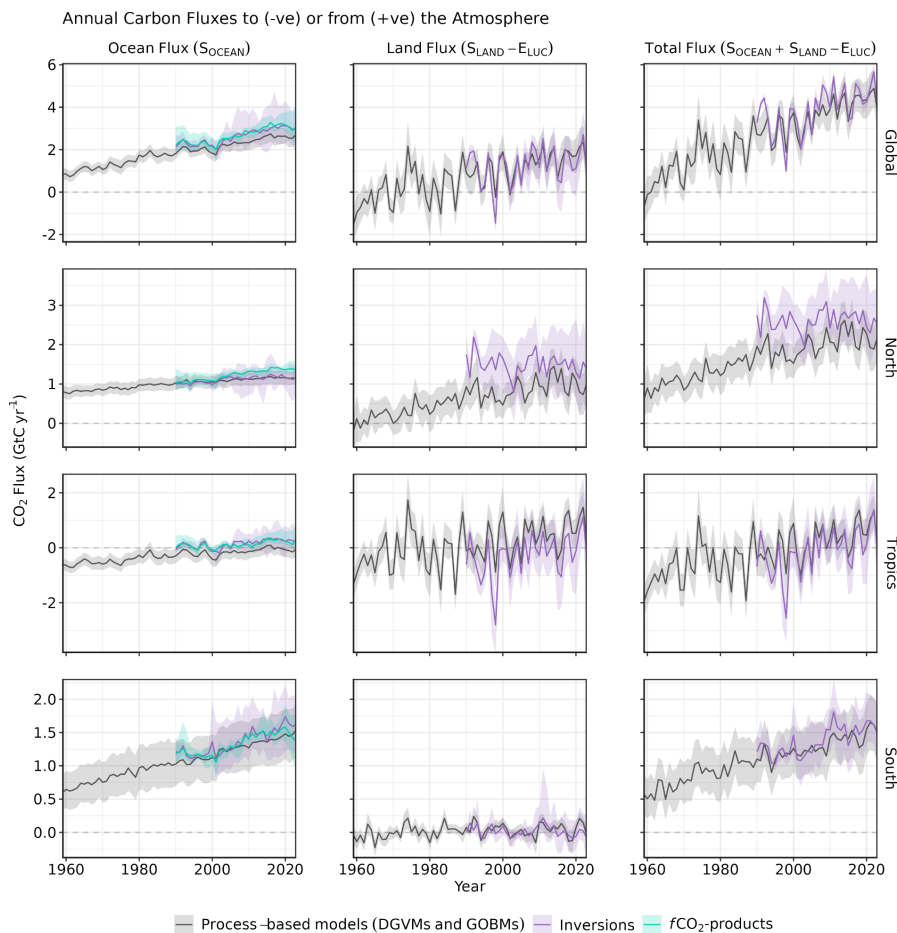
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**Figure 14.** CO<sub>2</sub> fluxes between the atmosphere and the Earth's surface separated between land and oceans, globally and in three latitude bands. The ocean flux is  $S_{\text{OCEAN}}$  and the land flux is the net atmosphere-land fluxes from the DGVMs. The latitude bands are (top row) global, (2<sup>nd</sup> row) north ( $>30^{\circ}\text{N}$ ), (3<sup>rd</sup> row) tropics ( $30^{\circ}\text{S}$ - $30^{\circ}\text{N}$ ), and (bottom row) south ( $<30^{\circ}\text{S}$ ), and over ocean (left column), land (middle column), and total (right column). Estimates are shown for: process-based models (DGVMs for land, GOBMs for oceans); inversion systems (land and ocean); and  $f\text{CO}_2$ -products (ocean only). Positive values are CO<sub>2</sub> sinks. Mean estimates from the combination of the process models for the land and oceans are shown (black line) with  $\pm 1\sigma$  of the model ensemble (grey shading). For the total uncertainty in the process-based estimate of the total sink, uncertainties are summed in quadrature. Mean estimates from the atmospheric inversions are shown (purple lines) with their full spread (purple shading). Mean estimates from the  $f\text{CO}_2$ -products are shown for the ocean domain (light blue

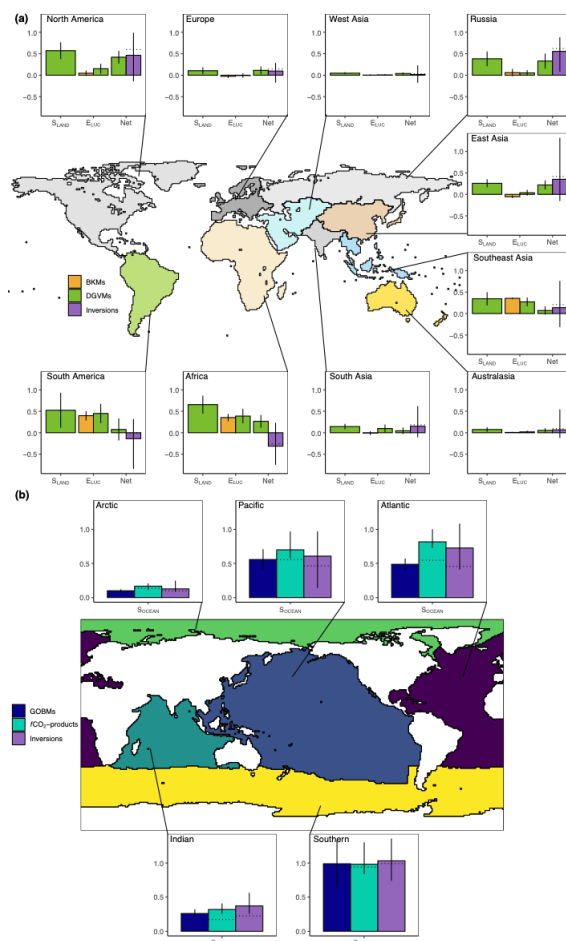
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lines) with full model spread (light blue shading). The global  $\text{SOCEAN}$  (upper left) and the sum of  $\text{SOCEAN}$  in all three regions represents the anthropogenic atmosphere-to-ocean flux based on the assumption that the preindustrial ocean sink was  $0 \text{ GtC yr}^{-1}$  when riverine fluxes are not considered. This assumption does not hold at the regional level, where preindustrial fluxes can be significantly different from zero. Hence, the regional panels for  $\text{SOCEAN}$  represent a combination of natural and anthropogenic fluxes. Bias-correction and area-weighting were only applied to global  $\text{SOCEAN}$ ; hence the sum of the regions is slightly different from the global estimate ( $<0.07 \text{ GtC yr}^{-1}$ ).

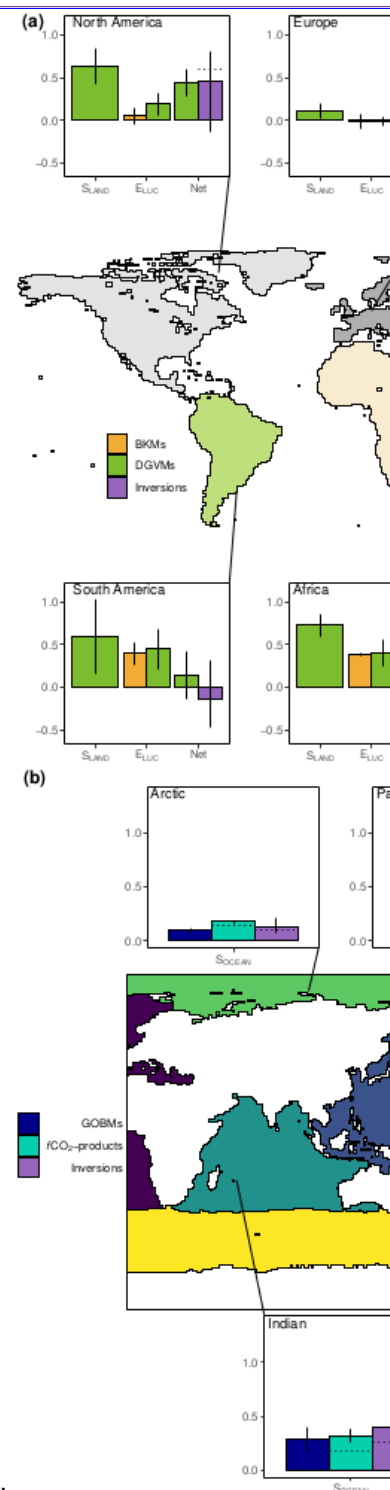
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**Figure 15.** Decadal mean (a) land and (b) ocean fluxes for RECCAP-2 regions over 2014-2023. For land fluxes,  $S_{LAND}$  is estimated by the DGVMs (green bars), with the error bar as  $\pm 1\sigma$  spread among models. A positive  $S_{LAND}$  is a net transfer of carbon from the atmosphere to the land.  $E_{LUC}$  fluxes are shown for both DGVMs (green) and bookkeeping models (orange), again with the uncertainty calculated as the  $\pm 1\sigma$  spread. Note, a positive  $E_{LUC}$  flux indicates a loss of carbon from the land. The net land flux is shown for both DGVMs (green) and atmospheric inversions (purple), including the full model spread for inversions. The net ocean sink ( $S_{OCEAN}$ ) is estimated by GOBMs (royal blue),  $fCO_2$ -products (cyan), and atmospheric inversions (purple). Uncertainty is estimated as the  $\pm 1\sigma$  spread for GOBMs, and the full model spread for the other two datasets. The dotted lines show the  $fCO_2$ -products and inversion results without river flux adjustment. Positive values are  $CO_2$  sinks.

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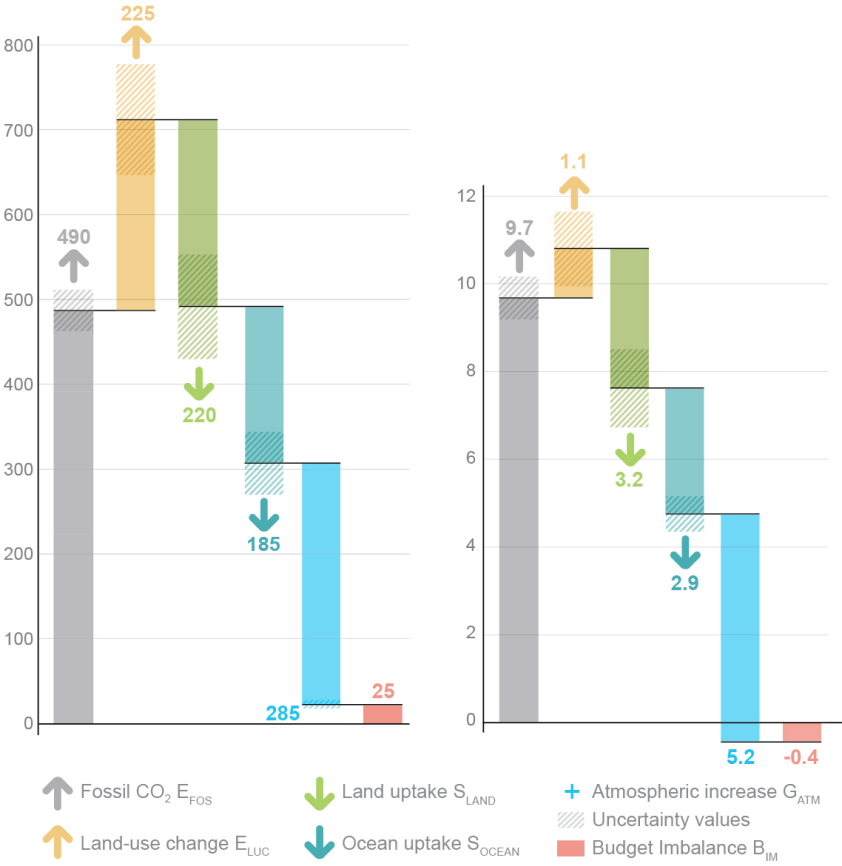
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# Anthropogenic carbon flows

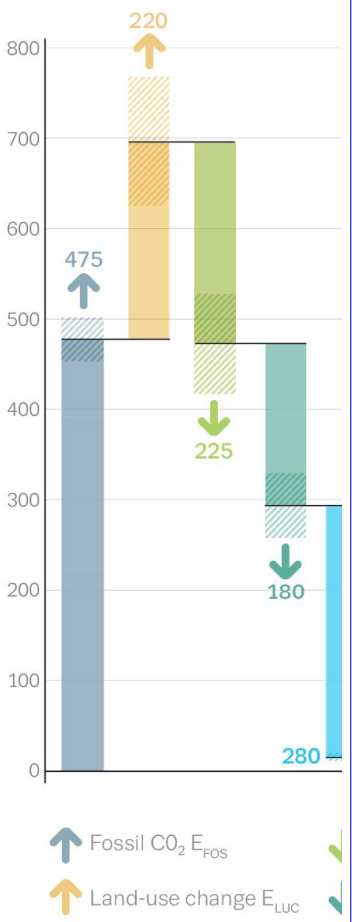
(a) Cumulative changes 1850-2023 GtC      (b) Mean fluxes 2014-2023 GtC per year



**Figure 16.** Cumulative changes over the 1850-2023 period (left) and average fluxes over the 2014-2023 period (right) for the anthropogenic perturbation of the global carbon cycle. See the caption of Figure 3 for key information and the methods in text for full details.

# Anthropogenic carbon flows

Cumulative changes 1850-2022 GtC

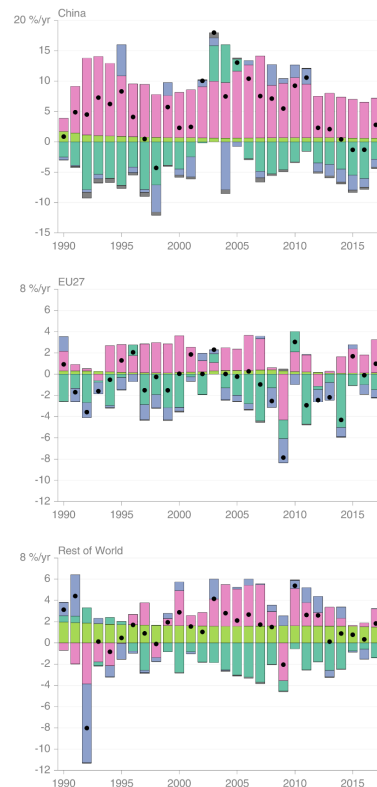


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**Figure 17.** Kaya decomposition of the main drivers of fossil CO<sub>2</sub> emissions, considering population, GDP per person, Energy per GDP, and CO<sub>2</sub> emissions per energy, for China (top left), USA (top right), EU27 (middle left), India (middle right), Rest of the World (bottom left), and World (bottom right). Black dots are the annual fossil CO<sub>2</sub> emissions growth rate, coloured bars are the contributions from the different drivers to this growth rate. A general trend is that population and GDP growth put upward pressure on emissions (positive values), while energy per GDP and, more recently, CO<sub>2</sub> emissions per energy put downward pressure on emissions (negative values). Both the COVID-19 induced drop during 2020 and the recovery in 2021 led to a stark contrast to previous years, with different drivers in each region.

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# 1 Global Carbon Budget [2024](#)

## 2 Supplementary Information

### 4 S.1 Methodology Fossil Fuel CO<sub>2</sub> emissions ([EFOS](#))

#### 5 S.1.1 Cement carbonation

6 From the moment it is created, cement begins to absorb CO<sub>2</sub> from the atmosphere, a process known as ‘cement  
7 carbonation’. We estimate this CO<sub>2</sub> sink, from 1931 onwards, as the average of two studies in the literature (Cao  
8 et al., 2020; Guo et al., 2021 extended by Huang et al., 2023). The Global Cement and Concrete Association  
9 reports a much lower carbonation rate, but this is based on the highly conservative assumption of 0% mortar  
10 (GCCA, 2021). Modelling cement carbonation requires estimation of a large number of parameters, including  
11 the different types of cement material in different countries, the lifetime of the structures before demolition, of  
12 cement waste after demolition, and the volumetric properties of structures, among others (Xi et al., 2016).  
13 Lifetime is an important parameter because demolition results in the exposure of new surfaces to the  
14 carbonation process. The main reasons for differences between the two studies appear to be the assumed  
15 lifetimes of cement structures and the geographic resolution, but the uncertainty bounds of the two studies  
16 overlap.

#### 17 S.1.2 Emissions embodied in goods and services

18 CDIAC, UNFCCC, and BP national emission statistics ‘include greenhouse gas emissions and removals taking  
19 place within national territory and offshore areas over which the country has jurisdiction’ (Rypdal et al., 2006),  
20 and are called territorial emission inventories. Consumption-based emission inventories allocate emissions to  
21 products that are consumed within a country, and are conceptually calculated as the territorial emissions minus  
22 the ‘embodied’ territorial emissions to produce exported products plus the emissions in other countries to  
23 produce imported products (Consumption = Territorial – Exports + Imports). Consumption-based emission  
24 attribution results (e.g. Davis and Caldeira, 2010) provide additional information to territorial-based emissions  
25 that can be used to understand emission drivers (Hertwich and Peters, 2009) and quantify emission transfers by  
26 the trade of products between countries (Peters et al., 2011a). The consumption-based emissions have the same  
27 global total, but reflect the trade-driven movement of emissions across the Earth’s surface in response to human  
28 activities. We estimate consumption-based emissions from 1990-2020 by enumerating the global supply chain  
29 using a global model of the economic relationships between economic sectors within and between every country  
30 (Andrew and Peters, 2013; Peters et al., 2011b). Our analysis is based on the economic and trade data from the  
31 Global Trade and Analysis Project (GTAP; Narayanan et al., 2015), and we make detailed estimates for the  
32 years 1997 (GTAP version 5), 2001 (GTAP6), and 2004, 2007, 2011, and 2014 (GTAP10.0a), covering 57  
33 sectors and 141 countries and regions. The detailed results are then extended into an annual time series from  
34 1990 to the latest year of the Gross Domestic Product (GDP) data (2020 in this budget), using GDP data by  
35 expenditure in current exchange rate of US dollars (USD; from the UN National Accounts main [Aggregates](#)  
36 database; UN, 2022) and time series of trade data from GTAP (based on the methodology in Peters et al.,

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2011b). We estimate the sector-level CO<sub>2</sub> emissions using the GTAP data and methodology, add the flaring and cement emissions from our fossil CO<sub>2</sub> dataset, and then scale the national totals (excluding bunker fuels) to match the emission estimates from the carbon budget. We do not provide a separate uncertainty estimate for the consumption-based emissions, but based on model comparisons and sensitivity analysis, they are unlikely to be significantly different than for the territorial emission estimates (Peters et al., 2012b).

### S.1.3 Uncertainty assessment for $E_{FOS}$

We estimate the uncertainty of the global fossil CO<sub>2</sub> emissions at  $\pm 5\%$  (scaled down from the published  $\pm 10\%$  at  $\pm 2\sigma$  to the use of  $\pm 1\sigma$  bounds reported here; Andres et al., 2012). This is consistent with a more detailed analysis of uncertainty of  $\pm 8.4\%$  at  $\pm 2\sigma$  (Andres et al., 2014) and at the high-end of the range of  $\pm 5\text{--}10\%$  at  $\pm 2\sigma$  reported by (Ballantyne et al., 2015). This includes an assessment of uncertainties in the amounts of fuel consumed, the carbon and heat contents of fuels, and the combustion efficiency. While we consider a fixed uncertainty of  $\pm 5\%$  for all years, the uncertainty as a percentage of emissions is growing with time because of the larger share of global emissions from emerging economies and developing countries (Marland et al., 2009). Generally, emissions from mature economies with good statistical processes have an uncertainty of only a few per cent (Marland, 2008), while emissions from strongly developing economies such as China have uncertainties of around  $\pm 10\%$  (for  $\pm 1\sigma$ ; Gregg et al., 2008; Andres et al., 2014). Uncertainties of emissions are likely to be mainly systematic errors related to underlying biases of energy statistics and to the accounting method used by each country.

### S.1.4 Growth rate in emissions

We report the annual growth rate in emissions for adjacent years (in percent per year) by calculating the difference between the two years and then normalising to the emissions in the first year:  $(E_{FOS}(t_0+1) - E_{FOS}(t_0))/E_{FOS}(t_0) \times 100\%$ . We apply a leap-year adjustment where relevant to ensure valid interpretations of annual growth rates. This affects the growth rate by about  $0.3\%$  yr<sup>-1</sup> ( $1/366$ ) and causes calculated growth rates to go up approximately  $0.3\%$  if the first year is a leap year and down  $0.3\%$  if the second year is a leap year.

The relative growth rate of  $E_{FOS}$  over time periods of greater than one year can be rewritten using its logarithm equivalent as follows:

$$\frac{1}{E_{FOS}} \frac{dE_{FOS}}{dt} = \frac{d(\ln E_{FOS})}{dt} \quad (2)$$

Here we calculate relative growth rates in emissions for multi-year periods (e.g. a decade) by fitting a linear trend to  $\ln(E_{FOS})$  in Eq. (2), reported in percent per year.

### S.1.5 Emissions projection for 2023

To gain insight on emission trends for 2023, we provide an assessment of global fossil CO<sub>2</sub> emissions,  $E_{FOS}$ , by combining individual assessments of emissions for China, USA, the EU, and India (the four countries/regions with the largest emissions), and the rest of the world.

The methods are specific to each country or region, as described in detail below.

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73 **China:** We use a regression between monthly data for each fossil fuel and cement, and annual data for  
74 consumption of fossil fuels / production of cement to project full-year growth in fossil fuel consumption and  
75 cement production. The monthly data for each product consists of the following:

- 76 · Coal: Production data from the National Bureau of Statistics (NBS), plus net imports from the China  
77 Customs Administration (i.e., gross supply of coal, not including inventory changes), adjusted  
78 using monthly production data for thermal electricity, crude steel, pig iron, coke and cement from  
79 NBS.
- 80 · Oil: Production data from NBS, plus net imports from the China Customs Administration (i.e., gross  
81 supply of oil, not including inventory changes)
- 82 · Natural gas: Same as for oil
- 83 · Cement: Production data from NBS

84 For oil, we use data for production and net imports of refined oil products rather than crude oil. This choice is  
85 made because refined products are one step closer to actual consumption, and because crude oil can be subject  
86 to large market-driven and strategic inventory changes that are not captured by available monthly data.  
87 Furthermore, refinery output in 2022 was atypically low through August of that year compared to the rest of the  
88 year, which results in very high growth figures for the 2023 data compared to what one can likely expect for the  
89 last four months of this year. The estimate has been adjusted down by 0.8 percentage points to account for this,  
90 corresponding to how much lower the ratio of January-August and September-December refinery output was in  
91 2022 compared to the average for 2014-2022.

92 For each fuel and cement, we make a Bayesian linear regression between year-on-year cumulative growth in  
93 supply (production for cement) and full-year growth in consumption (production for cement) from annual  
94 consumption data. In the regression model, the growth rate in annual consumption (production for cement) is  
95 modelled as a regression parameter multiplied by the cumulative year-on-year growth rate from the monthly  
96 data through August of each year for past years (through 2022). We use broad Gaussian distributions centered  
97 around 1 as priors for the ratios between annual and through-August growth rates. We then use the posteriors for  
98 the growth rates together with cumulative monthly supply/production data through August of 2023 to produce a  
99 posterior predictive distribution for the full-year growth rate for fossil fuel consumption / cement production in  
100 2023.

101 If the growth in supply/production through August were an unbiased estimate of the full-year growth in  
102 consumption/production, the posterior distribution for the ratio between the monthly and annual growth rates  
103 would be centered around 1. However, in practice the ratios are different from 1 (in most cases below 1). This is  
104 a result of various biasing factors such as uneven evolution in the first and second half of each year, inventory  
105 changes that are somewhat anti-correlated with production and net imports, differences in statistical coverage,  
106 and other factors that are not captured in the monthly data.

107 For fossil fuels, the mean of the posterior distribution is used as the central estimate for the growth rate in 2023,  
108 while the edges of a 68% credible interval (analogous to a 1-sigma confidence interval) are used for the upper  
109 and lower bounds.

110 **USA:** We use emissions estimated by the U.S. Energy Information Administration (EIA) in their Short-Term  
 111 Energy Outlook (STEO) for emissions from fossil fuels to get both YTD and a full year projection (EIA, 2023).  
 112 The STEO also includes a near-term forecast based on an energy forecasting model which is updated monthly  
 113 (we use the November 2023 edition), and takes into account expected temperatures, household expenditures by  
 114 fuel type, energy markets, policies, and other effects. We combine this with our estimate of emissions from  
 115 cement production using the monthly U.S. cement clinker production data from USGS for January-August  
 116 2023, assuming changes in clinker production over the first part of the year apply throughout the year.

117 **India:** We use monthly emissions estimates for India updated from Andrew (2020b) through August-October  
 118 2023. These estimates are derived from many official monthly energy and other activity data sources to produce  
 119 direct estimates of national CO<sub>2</sub> emissions, without the use of proxies. Emissions from coal are then extended to  
 120 October using a regression relationship based on power generated from coal, coal dispatches by Coal India Ltd.,  
 121 the composite PMI, time, and days per month. For the last 3-5 months of the year, each series is extrapolated  
 122 assuming typical (pre-2019) trends.

123 **EU:** We use a refinement to the methods presented by Andrew (2021), deriving emissions from monthly energy  
 124 data reported by Eurostat. Some data gaps are filled using data from the Joint Organisations Data Initiative  
 125 (JODI, 2022). Sub-annual cement and cement-clinker production data are limited, but data for Germany, Poland  
 126 and Spain, the three largest producers, suggest a decline of over 8%. For fossil fuels this provides estimates  
 127 through July-September, varying by fuel. We extend coal emissions through October using a regression model  
 128 built from generation of power from hard coal, power from brown coal, and the number of working days in  
 129 Germany, the biggest coal consumer in the EU. These are then extended through the end of the year assuming  
 130 typical trends. We extend oil emissions by building a regression model between our monthly CO<sub>2</sub> estimates and  
 131 oil consumption reported by the EIA for Europe in its Short-Term Energy Outlook (November edition), and then  
 132 using this model with EIA's monthly forecasts. For natural gas, the strong seasonal signal allows the use of the  
 133 bias-adjusted Holt-Winters exponential smoothing method (Chatfield, 1978), although this comes with larger  
 134 uncertainty given the unusual energy situation in Europe in 2022-23.

135 **Rest of the world:** We use the close relationship between the growth in GDP and the growth in emissions  
 136 (Raupach et al., 2007) to project emissions for the current year. This is based on a simplified Kaya Identity,  
 137 whereby  $E_{FOS}$  (GtC yr<sup>-1</sup>) is decomposed by the product of GDP (USD yr<sup>-1</sup>) and the fossil fuel carbon intensity of  
 138 the economy ( $I_{FOS}$ ; GtC USD<sup>-1</sup>) as follows:

$$139 \quad E_{FOS} = GDP \times I_{FOS} \quad (3)$$

140 Taking a time derivative of Equation (3) and rearranging gives:

$$141 \quad \frac{1}{E_{FOS}} \frac{dE_{FOS}}{dt} = \frac{1}{GDP} \frac{dGDP}{dt} + \frac{1}{I_{FOS}} \frac{dI_{FOS}}{dt} \quad (4)$$

142 where the left-hand term is the relative growth rate of  $E_{FOS}$ , and the right-hand terms are the relative growth  
 143 rates of GDP and  $I_{FOS}$ , respectively, which can simply be added linearly to give the overall growth rate.

144 The  $I_{FOS}$  is based on GDP in constant PPP (Purchasing Power Parity) from the International Energy Agency  
 145 (IEA) up to 2017 (IEA/OECD, 2019) and extended using the International Monetary Fund (IMF) growth rates  
 146 through 2022 (IMF, 2023). Interannual variability in  $I_{FOS}$  is the largest source of uncertainty in the GDP-based



emissions projections. We thus use the standard deviation of the annual IFOS for the period 2013-2022 as a measure of uncertainty, reflecting a  $\pm 1\sigma$  as in the rest of the carbon budget. For rest-of-world oil emissions growth, we use the global oil demand forecast published by the EIA less our projections for the other four regions, and estimate uncertainty as the maximum absolute difference over the period available for such forecasts using the specific monthly edition (e.g. August) compared to the first estimate based on more solid data in the following year (April).

**Bunkers:** Given the divergence in behaviour of international shipping from countries' emissions since the COVID-19 pandemic, we project international bunkers separately using sub-annual data on international aviation from the OECD (Clarke et al., 2022) and international shipping from MarineBenchmark and IMF (Cerdeiro et al., 2020).

**World:** The global total is the sum of each of the countries and regions.

## S.2 Methodology CO<sub>2</sub> emissions from land-use, land-use change and forestry (ELUC)

The net CO<sub>2</sub> flux from land-use, land-use change and forestry (ELUC, called land-use change emissions in the rest of the text) includes CO<sub>2</sub> fluxes from deforestation, afforestation, logging and forest degradation (including harvest activity), shifting cultivation (cycle of cutting forest for agriculture, then abandoning), regrowth of forests following wood harvest or abandonment of agriculture, peat burning, and peat drainage. Land-management activities are only partly included in our land-use change emissions estimates (Table S1). Some land-use change and land-management activities cause emissions of CO<sub>2</sub> to the atmosphere, while others remove CO<sub>2</sub> from the atmosphere. ELUC is the net sum of emissions and removals due to all anthropogenic activities considered. Our annual estimates for 1960-2022 are provided as the average of results from four bookkeeping approaches (Supplement S.2.1 below): the Bookkeeping of Land Use Emissions model (BLUE; Hansis et al., 2015), the compact Earth system model OSCAR (Gasser et al., 2020), an estimate from Houghton and Castanho (2023; hereafter H&C2023), and the Land-Use Change Emissions model (LUCE; Qin et al., 2024). Peat emissions are added from external datasets (see Supplement S.2.1 below). BLUE and OSCAR are updated with new land-use forcing data covering the time period until 2023. All four data sets are extrapolated to provide a projection for 2024 (see Supplement S.2.5 below). In addition, we use results from Dynamic Global Vegetation Models (DGVMS; see Supplement S.2.2 and Table 4) to help quantify the uncertainty in ELUC (Supplement S.2.4), and thus better characterise the robustness of annual estimates and trends. In this budget, we follow the scientific ELUC definition as used by global carbon cycle models, which counts fluxes due to environmental changes on managed land towards S<sub>LAND</sub>, as opposed to the national greenhouse gas inventories (NGHGs) under the UNFCCC, most of which include them in ELUC and thus often report smaller land-use emissions (Grassi et al., 2018; Petrescu et al., 2020). Following the methodology of Grassi et al. (2023), we provide harmonised estimates of the two approaches further below (see Supplement S.2.3).

### S.2.1 Bookkeeping models

CO<sub>2</sub> emissions and removals from land-use change are calculated by four bookkeeping models. These are based on the original bookkeeping approach of Houghton (2003), which keeps track of the carbon stored in vegetation and soils before and after a land-use change event (transitions between various natural vegetation types,

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croplands, and pastures). Literature-based response curves describe [the](#) decay of vegetation and soil carbon, including [carbon](#) transfer to product pools of different lifetimes, as well as carbon uptake due to regrowth. In addition, the bookkeeping models represent long-term degradation of primary forest as lowered standing vegetation and soil carbon stocks in secondary forests, and include forest management practices such as wood harvests.

BLUE, [LUCE](#) and H&C2023 exclude the transient response of land ecosystems to changes in climate, atmospheric CO<sub>2</sub>, and other environmental factors, and base the carbon densities of soil and vegetation on contemporary data from literature and inventory data. Since carbon densities thus remain fixed over time, the additional sink capacity that ecosystems provide in response to CO<sub>2</sub> fertilisation and some other environmental changes are not captured by these models (Pongratz et al., 2014). [OSCAR includes this transient response, and it follows a theoretical framework \(Gasser and Ciais, 2013\) that allows separating bookkeeping land-use emissions and the loss of additional sink capacity. Only the former is included here, while the latter is discussed in Supplement S.6.4. The bookkeeping models differ in \(1\) computational units \(spatially explicit treatment of land-use change at 0.25° resolution for BLUE and \[LUCE\]\(#\), country-level for H&C2023 and OSCAR\), \(2\) processes represented \(see Table S1\), and \(3\) carbon densities assigned to vegetation and soils for different types of vegetation \(literature-based for BLUE and H&C2023, calibrated to DGVMs for OSCAR, \[mainly literature-based but additionally considering the impact of land cohort age on secondary land carbon stocks for LUCE\]\(#\)\). A notable difference between models exists with respect to the treatment of shifting cultivation: H&C2023 assumes that forest loss—derived from the Global Forest Resources Assessment \(FRA; FAO, 2020\)—in excess of increases in cropland and pastures—derived from FAOSTAT \(FAO, 2021\)—represents an increase in shifting cultivation. If the excess loss of forests in a year is negative, it is assumed that shifting cultivation is returned to forest. Historical areas in shifting cultivation are defined taking into account country-based estimates of areas in fallow in 1980 \(FAO/UNEP, 1981\) and expert opinion \(from Heinimann et al., 2017\). In contrast, BLUE, \[OSCAR\]\(#\), and \[LUCE\]\(#\) include subgrid-scale transitions between all vegetation types. Furthermore, H&C2023 assumes conversion of natural grasslands to pasture, while BLUE, \[OSCAR\]\(#\), and \[LUCE\]\(#\) allocate pasture transitions proportionally to all natural vegetation that exists in a grid-cell. This is one reason for generally higher emissions in BLUE and OSCAR. In this GCB, we split CO<sub>2</sub> emissions into emissions from permanent deforestation and from deforestation for shifting cultivation. Similarly, we separate the forest \(re-\)growth estimates into \(re-\)growth from af/reforestation and from regrowth associated with shifting cultivation. This distinction is insightful with regard to the levers on the reduction of net emissions: as deforestation for shifting cultivation is only temporary, the associated CO<sub>2</sub> emissions cannot easily be avoided without compromising the CO<sub>2</sub> removals from regrowth in shifting cultivation cycles. By contrast, permanent deforestation is typically not directly related to af/reforestation. Stopping deforestation for permanent agricultural expansion and increasing the forest area provide two independent \[levers for net emissions reduction\]\(#\). Bookkeeping models do not directly capture carbon emissions from the organic layers of drained peat soils nor from peat fires. Particularly the latter can create large emissions and interannual variability due to synergies of land-use and climate variability in equatorial Southeast Asia, \[especially during El-Niño events\]\(#\). \[We\]\(#\) add peat fire emissions based on the Global Fire Emission Database \(GFED4s; van der Werf et al., 2017\) to the bookkeeping models' output. Peat fire emissions are calculated by multiplying the mass of dry matter emitted by peat fires with the C emission factor for peat fires indicated in the GFED4s database. Emissions from deforestation and](#)

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degradation fires used for extrapolating the H&C2023 data beyond 2020 and to derive the 2023 projection of all three models (see below) are calculated analogously. The satellite-derived GFED4s estimates of peat fire emissions start in 1997. For the previous years, we follow the approach by Houghton and Nassikas (2017), which linearly ramps up from zero emissions in 1980 to 0.04 GtC yr<sup>-1</sup> in 1996, reflecting the onset of major clearing of peatlands in equatorial Southeast Asia in the 1980s. We further add estimates of peat drainage emissions, combining estimates from three spatially explicit datasets. We employ FAO peat drainage emissions 1990–2022 from croplands and grasslands (Conchedda and Tubiello, 2020; FAO, 2023), peat drainage emissions 1700–2010 from simulations with the DGVM ORCHIDEE-PEAT (Qiu et al., 2021), and peat drainage emissions 1701–2023 from simulations with the DGVM LPX-Bern v1.5 (Lienert and Joos, 2018; Müller and Joos, 2021), the latter applying the updated LUH2 GCB2024 forcing as also used by BLUE, OSCAR, LUCE, and the DGVMs. The LPX-Bern simulations started from a transient run over the last deglaciation (20,050 to 1700 AD) following Müller and Joos (2020) and are forced by changes in climate, atmospheric CO<sub>2</sub>, nitrogen deposition/input, and land-use changes. Simulations were done with, and without prescribing land-use changes since 1700 AD. The difference between the simulations represents anthropogenic peat drainage emissions. To account for internal variability, we used the median peat drainage emissions from a 20-member ensemble. In LPX-Bern, peat carbon is stored in (i) active peatlands, (ii) former peatlands (“natural”), and (iii) former peatlands under anthropogenic use. We average the two CO<sub>2</sub> emission cases from Müller and Joos (2021), assuming that half the peat carbon is lost immediately to the atmosphere after transformation from active to former peatland, while the rest decays slowly, pending on local temperature and soil moisture. The LPX-Bern peat drainage emissions show a very high emission peak in Russia in 1959 followed by very low emissions in 1960. This peak can be attributed to an artefact in the HYDE3.4 dataset, which was corrected for Brazil and the Democratic Republic of the Congo in GCB2022 (Friedlingstein et al. 2022b) but remains for Russia where it strongly impacts the LPX-Bern peat drainage estimates in 1959 and 1960. To correct for this unrealistic peak, we replace the LPX-Bern peat drainage emissions in Russia in 1959 and 1960 by the average of the estimates in 1958 and 1961. FAO data are extrapolated to 1850–2023 by keeping the post-2020 emissions constant at 2020 levels and by linearly increasing tropical peat drainage emissions between 1980 and 1990 starting from 0 GtC yr<sup>-1</sup> in 1980 (consistent with H&N2017’s assumption, Houghton and Nassikas, 2017), and by keeping pre-1990 emissions from the often old drained areas of the extra-tropics constant at 1990 emission levels. ORCHIDEE-PEAT data are extrapolated to 2011–2023 by replicating the average emissions in 2000–2010 (pers. comm. C. Qiu). Further, ORCHIDEE-PEAT only provides peat drainage emissions north of 30°N, and thus we fill the regions south of 30°N by the average peat drainage emissions from FAO and LPX-Bern. The final peat drainage emissions are calculated as the average of the estimates from the three different peat drainage datasets. The net E<sub>LUC</sub> values indicated in the manuscript are the sum of E<sub>LUC</sub> estimates from bookkeeping models, peat fire emissions, and peat drainage emissions. The four bookkeeping estimates used in this study differ with respect to the land-use change data used to drive the models. H&C2023 base their estimates directly on the Forest Resource Assessment (FRA) of FAO, which provides statistics on forest-area change and management at intervals of five years currently updated until 2020 (FAO, 2020). The data is based on country reporting to FAO and may include remote-sensing information in more recent assessments. Changes in land use other than forests are based on annual, national changes in cropland and pasture areas reported by FAO (FAO, 2021). BLUE and LUCE use the harmonised land-use

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change data LUH2-GCB2024 covering the period 850-2023 (an update to the previously released LUH2 v2h dataset; Hurtt et al., 2017; Hurtt et al., 2020), which was also used as input to the DGVMs (Supplement S.2.2). LUH2-GCB2024 provides land-use change data at 0.25° spatial resolution based on the FAO data (as described in Supplement S.2.2) as well as the HYDE3.4 dataset (Klein Goldewijk et al., 2017a, 2017b), considering subgrid-scale transitions between primary forest, secondary forest, primary non-forest, secondary non-forest, cropland, pasture, rangeland, and urban land (Hurtt et al., 2020; Chini et al., 2021). LUH2-GCB2024 provides a distinction between rangelands and pasture, based on inputs from HYDE. Rangeland establishment in forests is assumed to transform forests to grasslands, rangeland establishment in non-forest primary vegetation degrades primary to secondary vegetation, and rangeland establishment in non-forest secondary vegetation has no effect (e.g., browsing on shrubland) (Ma et al., 2020). This case distinction is implemented in BLUE, based on a forest mask provided with LUH2-GCB2021. OSCAR was run with both LUH2-GCB2024 and FAO/FRA, where the drivers of the latter were linearly extrapolated to 2023 using their 2015-2020 trends. The best-guess OSCAR estimate used in our study is a combination of results for LUH2-GCB2024 and FAO/FRA land-use data and a large number of perturbed parameter simulations weighted against a constraint (the cumulative  $S_{LAND}$  over 1960-2022 of last year's GCB). As the record of H&C2023 ends in 2020, we extend it up to 2023 by adding the yearly anomalies of the emissions from tropical deforestation and degradation fires from GFED4s between 2020 and 2022 to the model's estimate for 2020 (emissions from peat fires and peat drainage are added to all models later in the process).

The annual  $E_{LUC}$  from 1850 onwards is calculated as the average of the estimates from BLUE, H&C2023, OSCAR, and LUCE. For the cumulative numbers starting in 1750, emission estimates between 1750-1850 are added based on the average of four earlier publications ( $30 \pm 20$  GtC 1750-1850, rounded to nearest 5; Le Quéré et al., 2016).

We provide a split of net  $E_{LUC}$  into component fluxes to better identify reasons for divergence between bookkeeping estimates and to give more insight into the drivers of net  $E_{LUC}$ . This split distinguishes between emissions from deforestation (including due to shifting cultivation), removals from forest (re-)growth (including regrowth in shifting cultivation cycles), fluxes from wood harvest and other forest management (i.e., emissions in forests from slash decay and emissions from product decay following wood harvesting, removals from regrowth after wood harvesting, and fire suppression), emissions from peat drainage and peat fires, and emissions and removals associated with all other land-use transitions. Additionally, we split deforestation emissions into emissions from permanent deforestation and emissions from deforestation in shifting cultivation cycles, and we split removals from forest (re-)growth into forest (re-)growth due to afforestation and forest regrowth in shifting cultivation cycles. This split helps to identify the emission reductions that would be achievable by halting permanent deforestation, and the removals that are caused by permanently increasing the forest cover through re/afforestation. Forest (re-)growth due to afforestation is calculated using a slightly updated method compared to GCB2023, now following the method used to calculate CDR due to re/afforestation in the 2nd State of CDR Report (Pongratz et al., 2024).  $E_{LUC}$  data are provided as global sums, as spatially explicit estimates at 0.25° spatial resolution (i.e., the native LUH2 resolution), and for 199 countries (based on the list of UNFCCC parties). Spatially explicit  $E_{LUC}$  estimates for BLUE and LUCE are directly available at 0.25°. For OSCAR and H&C2023, the country-level estimates were scaled to 0.25° based on the patterns of gross emissions and gross removals in BLUE (see Schwingshackl et al. 2022 for more details about

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the methodology). The gridded net  $E_{LUC}$  estimates of BLUE, LUCE, OSCAR, and H&C2023 are averaged, and the gridded estimates of peat drainage emissions (average of FAO, LPX-Bern, and ORCHIDEE-PEAT) and of peat fire emissions (from GFED4s) are added. Country-level estimates for the gridded datasets (BLUE, LUCE, LPX-Bern, ORCHIDEE-PEAT, GFED4s) are calculated based on a country map from Eurostat (Eurostat, 2024), which was remapped to 0.25°. In case multiple countries are present in a 0.25° grid cell, the  $E_{LUC}$  estimates are allocated proportional to each country's land fraction in that grid cell.

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### 400 S.2.2 Dynamic Global Vegetation Models (DGVMs)

Land-use change CO<sub>2</sub> emissions are also estimated by an ensemble of 20 DGVMs. The DGVMs account for deforestation and regrowth, the most important components of  $E_{LUC}$ , but they do not represent all processes resulting directly from human activities on land (Table S1). All DGVMs represent processes of vegetation growth and mortality, as well as decomposition of dead organic matter associated with natural cycles, and include the vegetation and soil carbon response to increasing atmospheric CO<sub>2</sub> concentration, to climate variability and to climate change. Most models explicitly simulate the coupling of carbon and nitrogen cycles and account for atmospheric N deposition and N fertilisers (Table S1). The DGVMs are independent from the other budget terms except for their use of atmospheric CO<sub>2</sub> concentration to calculate the fertilisation effect of CO<sub>2</sub> on plant photosynthesis.

All DGVMs use the LUH2-GCB2024 dataset as input, which includes the HYDE cropland/grazing land dataset (Klein Goldewijk et al., 2017a, 2017b), and some additional information on land-use transitions, land-use management activities and wood harvest. This includes annual, quarter-degree (regridged from 5 minute resolution), fractional data on cropland and pasture from HYDE3.4.

DGVMs that do not simulate subgrid-scale transitions (i.e., those estimating net land-use emissions; see Table S1) used the HYDE information on agricultural area change. For all countries, with the exception of Brazil, the Democratic Republic of the Congo, Indonesia, and China these data are based on the available annual FAO statistics of change in agricultural land area available from 1961 up to and including 2017. The FAO retrospectively revised their reporting for the Democratic Republic of the Congo, which was newly available until 2020 as reported in GCB2022. In addition to FAO country-level statistics, the HYDE3.4 cropland/grazing land dataset is constrained spatially based on multi-year satellite land cover maps from ESA CCI LC (see below). The extension of HYDE beyond the years that were directly informed by data was done as part of the LUH2 methodology this year and was a simple extension of the previous 5-year trend. The actual years for this extension varied by country since some countries were based on FAO data (2021), some used the China data (2019), and some used MapBiomass data (Brazil and Indonesia, 2022). This methodology is not appropriate for countries that have experienced recent rapid changes in the rate of land-use change, e.g. Brazil which has experienced a recent upturn in deforestation. For Brazil and Indonesia we replace FAO state-level data for cropland and grazing land in HYDE by those from the satellite-based land cover dataset MapBiomass (collection 7) for 1985-2022 (Brazil) (Souza et al. 2020) and 2000-2022 (Indonesia). ESA-CCI is used to spatially disaggregate as described below. The pre-1985 period is scaled with the per capita numbers from 1985 from MapBiomass, so this transition is smooth.

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450 HYDE uses satellite imagery from ESA-CCI from 1992–2018 for more detailed yearly allocation of cropland  
 451 and grazing land, with the ESA area data scaled to match the FAO annual totals at country-level. The original  
 452 300 metre spatial resolution data from ESA was aggregated to a 5 arc minute resolution according to the  
 453 classification scheme as described in Klein Goldewijk et al. (2017a).  
 454 DGVMs that simulate subgrid-scale transitions (i.e., [those estimating](#) gross land-use emissions; see Table S1)  
 455 use more detailed land use transition and wood harvest information from the LUH2-[GCB2024](#) data set. LUH2-  
 456 [GCB2024](#) is an update of the comprehensive harmonised land-use data set (Hurtt et al., 2020), that includes  
 457 fractional data on primary and secondary forest vegetation, as well as all underlying transitions between land-  
 458 use states (850-[2023](#); Hurtt et al., 2011, 2017, 2020; Chini et al., 2021; Table S1). This data set consists of [0.25°](#)  
 459 fractional areas of land-use states and all transitions between those states, including a new wood harvest  
 460 reconstruction, new representation of shifting cultivation, crop rotations, management information including  
 461 irrigation and fertiliser application. The land-use states include five different crop types in addition to splitting  
 462 grazing land into managed pasture and rangeland. Wood harvest patterns are constrained with Landsat-based  
 463 tree cover loss data (Hansen et al. 2013). Updates of LUH2-[GCB2024](#) over last year's version (LUH2-  
 464 [GCB2023](#)) are using the most recent HYDE release, [HYDE4.3 is based on new FAO inputs for years 1961-](#)  
 465 [2021, new MapBiomass inputs for Brazil \(for years 1985-2022\) and Indonesia \(for years 2000-2022\) and new](#)  
 466 [cropland data for China from Yu et al. 2022 \(for years 1900-2019\).](#)  
 467 We use updated FAO wood harvest data for all dataset years from 1961 to [2022](#), and linearly extended to the  
 468 year 2023. The HYDE3 [4](#) population data is also used to extend the wood harvest time series back in time.  
 469 Other wood harvest inputs (for years prior to 1961) remain the same in LUH2. These updates in the land-use  
 470 forcing are shown in Figure [S7](#) in comparison to [LUH2-GCB2022](#) and [LUH2-GCB2023](#). DGVMs implement  
 471 land-use change [in different ways](#) (e.g. an increased cropland fraction in a grid cell can either be at the expense  
 472 of grassland, shrubs, or forest, the latter resulting in deforestation; land cover fractions of the non-agricultural  
 473 land differ between models). Similarly, model-specific assumptions are applied to convert deforested biomass or  
 474 deforested area, and other forest product pools into carbon, and different choices are made regarding the  
 475 allocation of rangelands as natural vegetation or pastures.  
 476 The difference between two DGVMs simulations (see Supplement S.4.1 below), one forced with historical  
 477 changes in land-use and a second one with time-invariant pre-industrial land cover and pre-industrial wood  
 478 harvest rates, allows quantification of the dynamic evolution of vegetation biomass and soil carbon pools in  
 479 response to land-use change in each model ( $E_{LUC}$ ). Using the difference between these two DGVM simulations  
 480 to diagnose  $E_{LUC}$  means the DGVM estimate includes the loss of additional sink capacity (around  $0.4 \pm 0.3$  GtC  
 481 yr<sup>-1</sup>; see Section 2.10 and Supplement S.6.4), while the bookkeeping model estimate does not.  
 482 As a criterion for inclusion in this carbon budget, we only retain models that simulate a positive  $E_{LUC}$  during the  
 483 1990s, as assessed in the IPCC AR4 (Denman et al., 2007) and AR5 (Ciais et al., 2013). All DGVMs met this  
 484 criterion.

### 486 S.2.3 Translation between [NGHGs](#) and $E_{LUC}$

487 [Land-use](#) emissions estimates from bookkeeping models and from national GHG Inventories ([NGHGs](#)) show a  
 488 [large gap](#) (see [Figure 8](#) and [Table S10](#)). This gap is due to different approaches for calculating “anthropogenic”

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CO<sub>2</sub> fluxes related to land-use change and land management (Grassi et al. 2018). Land sinks due to environmental change on managed lands are treated as non-anthropogenic in the global carbon budget, while they are generally considered as anthropogenic in NGHGs (“indirect anthropogenic fluxes”; Eggleston et al., 2006). Building on previous studies (Grassi et al. 2021), we implement an approach that adds the DGVM estimates of CO<sub>2</sub> fluxes due to environmental change from managed forest areas (part of S<sub>LAND</sub>) to the E<sub>LUC</sub> estimate from bookkeeping models. This sum is expected to be conceptually more comparable to NGHGI estimates than E<sub>LUC</sub>.

E<sub>LUC</sub> data are taken from bookkeeping models, in line with the global carbon budget approach. To determine S<sub>LAND</sub> in managed forest, the following steps were taken: Spatially gridded data of “natural” forest NBP (S<sub>LAND</sub> i.e., including carbon fluxes due to environmental change and excluding land use change fluxes) were obtained from DGVMs using S2 runs from the TRENDY v13 dataset. Results were first masked with a forest map that is based on tree cover data from Hansen et al. (2013). To perform the conversion “tree” cover to “forest” cover, we exclude gridcells with less than 20% tree cover and isolated pixels with maximum connectivity less than 0.5 ha following the FAO definition of forest. Forest NBP is then further masked with a map of “intact” forest for the year 2013, i.e. forest areas characterised by no remotely detected signs of human activity (Potapov et al. 2017). This way, we obtained S<sub>LAND</sub> in “intact” and “non-intact” forest areas, which previous studies (Grassi et al. 2021) indicated to be a good proxy, respectively, for “unmanaged” and “managed” forest areas in the NGHGI. Note that only a subset of models had forest NBP at grid cell level. For the other DGVMs, when a grid cell had forest, all the NBP in that grid cell was allocated to forest. Since S2 simulations use pre-industrial forest cover masks that are at least 20% larger than today’s forest (Hurt et al. 2020), we corrected this NBP by a ratio between observed (based on Hansen et al. 2013) and prescribed (from DGVMs) forest cover. This ratio is calculated for each individual DGVM that provides information on prescribed forest cover, and a common ratio (median ratio of this subset of models) is used. The details of the method used are explained in a GitHub repository (Alkama, 2022).

LULUCF data from NGHGs are from Grassi et al. (2023), updated up to August 2024. While Annex I countries report a complete time series 1990-2021, gap-filling was applied for Non-Annex I countries through linear interpolation between two points and/or through extrapolation backward (till 2000) and forward (till 2021) using the single closest available data. For all countries, the estimates of the years 2022 and 2023 are assumed to be equal to those of 2021. The managed forest area, used to filter S<sub>LAND</sub> data from DGVMs to derive the natural land sink in managed forests, accounts for temporal dynamics from 2000 to 2023. This data includes all CO<sub>2</sub> fluxes from land considered managed, which in principle encompasses all land uses (forest land, cropland, grassland, wetlands, settlements, and other land), changes among them, emissions from organic soils (i.e., from peat drainage) and from fires. In practice, although almost all Annex I countries report all land uses, many non-Annex I countries report only on deforestation and forest land, and only few countries report on other land uses. In most cases, NGHGs include most of the natural response to recent environmental change because they use direct observations (e.g., national forest inventories) that do not allow separating direct and indirect anthropogenic effects (Eggleston et al., 2006).

Last, we also used the gridded data of net land flux from 14 atmospheric inversion systems (Table S4) to get an additional estimate of land-use fluxes in managed land. We applied a correction for riverine transport (see Supplement S.5.1.) and multiplied the resulting values with the fraction of managed land in each grid cell for

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each inversion. For this purpose, we used masks of managed land from Grassi et al. (2023) available for the years 1994, 2002, 2010, and 2016. We linearly interpolated the masks in time and replicated the 2016 mask in the years 2017-2023. Subsequently, we applied another correction for lateral transport due to international wood and crop trade (data from Deng et al. 2024). The obtained values are summed globally and compared to the NGHGI estimates and the translated  $E_{LUC}$  estimates. Figure 8 and Table S10 shows the resulting translation of global carbon cycle models' land flux definitions to that of the NGHGI (discussed in Section 3.2.2). For comparison we also show LULUCF estimates from FAOSTAT (FAO, 2024), which include emissions from net forest conversion and fluxes on forest land (Tubiello et al., 2021) as well as CO<sub>2</sub> emissions from peat drainage and peat fires. Forest land stock change data for 2021-2023 are carried forward from the 2020 estimates. The FAO data shows global emissions of 0.30 GtC yr<sup>-1</sup> averaged over 2014-2023, in contrast to the removals of -0.76 GtC yr<sup>-1</sup> estimated by the gap-filled NGHGI data. Most of this difference is attributable to different scopes: a focus on carbon fluxes for the NGHGI and a focus on land-use area and biomass estimates for FAO. In particular, the NGHGI data includes a larger forest sink for non-Annex 1 countries resulting from a more complete coverage of non-biomass carbon pools and non-forest land uses. NGHGI and FAO data also differ in terms of underlying data on forest land (Grassi et al., 2022).

#### S.2.4 Uncertainty assessment for $E_{LUC}$

Differences between the bookkeeping models and DGVMs originate from three main sources: different methodologies, which among others lead to inclusion of the loss of additional sink capacity in DGVMs (see Supplement S.6.4), different underlying land-use/land cover datasets, and different processes represented (Table S1). We examine both the results from DGVMs and from the bookkeeping method and use the resulting variations as a way to characterise the uncertainty in  $E_{LUC}$ . Despite the existing differences, the  $E_{LUC}$  estimate from the DGVM multi-model mean is consistent with the average of the emissions from the bookkeeping models (Table 5). However, there are large differences among individual DGVMs (standard deviation at 0.6 GtC yr<sup>-1</sup>; Table 5), between the bookkeeping estimates (standard deviation at 0.3 GtC yr<sup>-1</sup> for cumulative emissions in 1850-2022), and between the H&C2023 model and its previous model version H&N2017 (average difference 1850-2015 of 0.2 GtC yr<sup>-1</sup>; see Table 1 in Houghton and Castanho, 2023). A factorial analysis of differences between BLUE and H&N2017 (the precursor of H&C2023) attributed them particularly to differences in carbon densities between primary and secondary vegetation (Bastos et al., 2021). Earlier studies additionally showed the relevance of the different land-use forcing as applied (in updated versions) also in the current study (Gasser et al., 2020). Ganzenmüller et al. (2022) showed that  $E_{LUC}$  estimates with BLUE are substantially smaller when the model is driven by a new high-resolution land-use dataset (HILDA+). They identified shifting cultivation and the way it is implemented in LUH2 as a main reason for this divergence. They further showed that a higher spatial resolution reduces the estimates of both gross emissions and gross removals because successive transitions are not adequately represented at coarser resolution, which has the effect that—despite capturing the same extent of transition areas—overall less area remains pristine at the coarser compared to the higher resolution.

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616 The uncertainty in  $E_{LUC}$  of  $\pm 0.7 \text{ GtC yr}^{-1}$  reflects our best value judgement that there is at least 68% chance  
 617 ( $\pm 1\sigma$ ) that the true land-use change emissions lie within the given range, for the range of processes considered  
 618 here. Prior to the year 1959, the uncertainty in  $E_{LUC}$  is taken from the standard deviation of the DGVMs. We  
 619 assign low confidence to the annual estimates of  $E_{LUC}$  because of the inconsistencies among estimates and  
 620 because of the difficulties to quantify some of the processes with DGVMs.

621

## 622 S.2.5 Land-use emissions projection for 2024

623 We project the 2024 land-use emissions for BLUE, H&C2023, OSCAR, and LUCE based on their  $E_{LUC}$   
 624 estimates for 2023 and on the interannual variability of peat fires and tropical deforestation and degradation fires  
 625 as estimated using active fire data (MCD14ML; Giglio et al., 2016). The latter scales almost linearly with GFED  
 626 emissions estimates over large areas (van der Werf et al., 2017), and thus allows for tracking fire emissions in  
 627 deforestation and tropical peat zones in near-real time. Peat drainage is assumed to be unaltered, as it has low  
 628 interannual variability. We project the 2024 land-use emissions for BLUE, H&C2023, OSCAR, and LUCE  
 629 based on their  $E_{LUC}$  estimates for 2023 and add the change in carbon emissions from peat fires and tropical  
 630 deforestation and degradation fires (2024 emissions relative to 2023 emissions) from GFED4s. The GFED4s  
 631 estimates for 2024 are as of October 17.

632

## 633 S.3 Methodology Ocean CO<sub>2</sub> sink

### 634 S.3.1 Observation-based estimates

635 We primarily use the observational constraints assessed by IPCC of a mean ocean CO<sub>2</sub> sink of  $2.2 \pm 0.7 \text{ GtC yr}^{-1}$   
 636 for the 1990s (90% confidence interval; Ciais et al., 2013) to verify that the GOBMs provide a realistic  
 637 assessment of SOCEAN. This is based on indirect observations with seven different methodologies and their  
 638 uncertainties, and further using three of these methods that are deemed most reliable for the assessment of this  
 639 quantity (Denman et al., 2007; Ciais et al., 2013). The observation-based estimates use the ocean/land CO<sub>2</sub> sink  
 640 partitioning from observed atmospheric CO<sub>2</sub> and O<sub>2</sub>/N<sub>2</sub> concentration trends (Manning and Keeling, 2006;  
 641 Keeling and Manning, 2014), an oceanic inversion method constrained by ocean biogeochemistry data  
 642 (Mikaloff Fletcher et al., 2006), and a method based on penetration time scale for chlorofluorocarbons (McNeil  
 643 et al., 2003). The IPCC estimate of  $2.2 \text{ GtC yr}^{-1}$  for the 1990s is consistent with a range of methods  
 644 (Wanninkhof et al., 2013). We refrain from using the IPCC estimates for the 2000s ( $2.3 \pm 0.7 \text{ GtC yr}^{-1}$ ), and the  
 645 period 2002-2011 ( $2.4 \pm 0.7 \text{ GtC yr}^{-1}$ , Ciais et al., 2013) as these are based on trends derived mainly from  
 646 models and one data-product (Ciais et al., 2013). Additional constraints summarised in AR6 (Canadell et al.,  
 647 2021) are the interior ocean anthropogenic carbon change (Gruber et al., 2019) and ocean sink estimate from  
 648 atmospheric CO<sub>2</sub> and O<sub>2</sub>/N<sub>2</sub> (Tohjima et al., 2019) which are used for model evaluation and discussion,  
 649 respectively.

650 We also use nine estimates of the ocean CO<sub>2</sub> sink and its variability based on surface ocean  $f\text{CO}_2$  maps obtained  
 651 by the interpolation of surface ocean  $f\text{CO}_2$  measurements. Seven of the methods cover a period from 1990  
 652 onwards due to severe restriction in data availability prior to 1990 (Figure 11), whereas two span the time period  
 653 from 1957 and 1959 onwards. These estimates differ in many respects: they use different maps of surface  $f\text{CO}_2$ ,

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different atmospheric CO<sub>2</sub> concentrations, wind products and different gas-exchange formulations as specified in Table S3. We refer to them as *f*CO<sub>2</sub>-products. The measurements underlying the surface *f*CO<sub>2</sub> maps are from the Surface Ocean CO<sub>2</sub> Atlas version 2024 (SOCAT v2024; Bakker et al., 2024), which is an update of version 3 (Bakker et al., 2016) and the subsequent annual updates used in previous versions of the global carbon budget. SOCAT v2024 has an additional 3.0 million *f*CO<sub>2</sub> measurements with an estimated accuracy of better than 5  $\mu$ atm relative to v2023. Of these, 2 million are from 2023 in a total of 210 data sets (Table S7), while the largest addition from earlier years is from 2022 with 64 data sets new to SOCAT. For the 2023 data, there are a total of 178 data sets with measurements in the Northern hemisphere, while there are only 52 with data from the Southern hemisphere. For the Southern Ocean, there are only 11 data sets from 2023 in the subpolar zone and further south (defined as south of 45°S), and only one from Austral winter (June-August). The coverage of SOCAT observations in 2023 is only about 50% of that in 2016 (Fig. 11), with large reductions in sampling in both the Northern (from 391 to 178 data sets) as well as Southern hemisphere (from 109 to 52 data sets). This reduction cannot be explained only in terms of lags in data submission. The quality control criteria used for SOCATv2024 are described in Lauvset et al. (2018). Each of the data-based estimates uses a different method to map the SOCAT v2024 data to the global ocean. The methods include a data-driven diagnostic method combined with a multi linear regression approach to extend back to 1957 (Rödenbeck et al., 2022; referred to here as Jena-MLS), four neural network models (Landschützer et al., 2014; referred to as VLIZ-SOMFFN; Chau et al., 2022; Copernicus Marine Environment Monitoring Service, referred to here as CMEMS-LSCE-FFNN; Zeng et al., 2022; referred to as NIES-ML3; Gregor et al. 2019, referred to as CSIR-ML6), one cluster regression approach (Gregor et al., 2024; referred to as OceanSODA-ETHZv2), a multi-linear regression method (Iida et al., 2021; referred to as JMA-MLR), and one method that relates the *f*CO<sub>2</sub> misfit between GOBMs and SOCAT to environmental predictors using the extreme gradient boosting method extending back to 1959 (Gloege et al., 2022). The ensemble mean of the *f*CO<sub>2</sub>-based flux estimates is calculated from these eight mapping methods. Further, we show the flux estimate of the UExp-FNN-U method (Watson et al., 2020; Ford et al., accepted), who also use a neural network model to map *f*CO<sub>2</sub> data to the globe, but resulting in a substantially larger ocean sink estimate, owing to a number of adjustments they applied to the surface ocean *f*CO<sub>2</sub> data. Concretely, these authors adjusted the SOCAT *f*CO<sub>2</sub> downward to account for differences in temperature between the depth of the ship intake and the relevant depth right near the surface, and included a further adjustment to account for the cool surface skin temperature effect. In Friedlingstein et al. 2023, the UExp-FNN-U product correction was applied illustrating that this temperature adjustment leads to an upward correction of the ocean carbon sink, up to 0.9 GtC yr<sup>-1</sup>, that, if correct, should be applied to all *f*CO<sub>2</sub>-based flux estimates. This year, the updated UExp-FFN-U method applies a smaller adjustment as proposed by Dong et al. (2022), who illustrate a smaller correction effect of 0.6 GtC yr<sup>-1</sup>. The impact of the cool skin effect on air-sea CO<sub>2</sub> flux is based on established understanding of temperature gradients (as discussed by Goddijn-Murphy et al., 2015 and Woolf et al., 2016), and laboratory observations (Jähne and Haussecker, 1998; Jähne, 2019), but in situ field observational evidence is lacking (Dong et al., 2022). The UExp-FNN-U method is thus, similar to the UExp-FNN-U flux estimate in previous editions, not included in the ensemble mean of the *f*CO<sub>2</sub>-based flux estimates. This choice will be re-evaluated in upcoming budgets based on further lines of evidence.

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**Deleted:** The Watson et al. flux estimate hence differs from the others by their choice of adjusting the flux to a cool, salty ocean surface skin. Watson et al. (2020) showed

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Typically,  $f\text{CO}_2$ -products do not cover the entire ocean due to missing coastal oceans and sea ice cover. The  $\text{CO}_2$  flux from each  $f\text{CO}_2$ -based product is already at or above 99% coverage (either due to complete coverage or a posteriori filling) of the ice-free ocean surface area in several products this year (UEXP-FNN-U, JMA-MLR, VLIZ-SOMFFN, Jena-MLS, OceanSODA-ETHZv2). The products that remained below 99% coverage of the ice-free ocean (CMEMS-LSCE-FFNN, NIES-ML3, UEXP-FNN-U, CSIR-ML6) were scaled by the following procedure:

Since v2022 of the GCB we now scale fluxes globally and regionally (North, Tropics, South) to match the ice-free area (using the HadISST sea surface temperature and sea ice cover; Rayner et al., 2003):

$$FCO_2^{reg-scaled} = \frac{A_{(1-ice)}^{region}}{A_{FCO_2}^{region}} \cdot FCO_2^{region}$$

In the equation,  $A$  represents area,  $(1 - ice)$  represents the ice free ocean,  $A_{FCO_2}^{region}$  represents the coverage of the  $f\text{CO}_2$ -product for a region, and  $FCO_2^{region}$  is the integrated flux for a region.

We further use results from two diagnostic ocean models, Khatiwala et al. (2013) and DeVries (2014), to estimate the anthropogenic carbon accumulated in the ocean prior to 1959. The two approaches assume constant ocean circulation and biological fluxes, with  $\text{SO}_{\text{OCEAN}}$  estimated as a response in the change in atmospheric  $\text{CO}_2$  concentration calibrated to observations. The uncertainty in cumulative uptake of  $\pm 20 \text{ GtC}$  (converted to  $\pm 1\sigma$ ) is taken directly from the IPCC's review of the literature (Rhein et al., 2013), or about  $\pm 30\%$  for the annual values (Khatiwala et al., 2009).

### S.3.2 Global Ocean Biogeochemistry Models (GOBMs)

The ocean  $\text{CO}_2$  sink for 1959-2023 is estimated using ten GOBMs (Table S2). The GOBMs represent the physical, chemical, and biological processes that influence the surface ocean concentration of  $\text{CO}_2$  and thus the air-sea  $\text{CO}_2$  flux. The GOBMs are forced by meteorological reanalysis and atmospheric  $\text{CO}_2$  concentration data available for the entire time period. They mostly differ in the source of the atmospheric forcing data (meteorological reanalysis), spin up strategies, and in their horizontal and vertical resolutions (Table S2). All GOBMs except one (CESM-ETHZ) do not include the effects of anthropogenic changes in nutrient supply (Duce et al., 2008). They also do not include the perturbation associated with changes in riverine organic carbon (see Section 2.10 and Supplement S.6.3).

Four sets of simulations were performed with each of the GOBMs. Simulation A applied historical changes in climate and atmospheric  $\text{CO}_2$  concentration. Simulation B is a control simulation with constant atmospheric forcing (normal year or repeated year forcing) and constant pre-industrial atmospheric  $\text{CO}_2$  concentration. Simulation C is forced with historical changes in atmospheric  $\text{CO}_2$  concentration, but repeated year or normal year atmospheric climate forcing. Simulation D is forced by historical changes in climate and constant pre-industrial atmospheric  $\text{CO}_2$  concentration.

The atmospheric  $\text{CO}_2$  forcing file was updated in GCB2024 to ensure consistency with the atmospheric  $\text{CO}_2$  growth rate reported in the GCB. Since January 1980, we use the  $\text{CO}_2$  global growth rate reported by NOAA/GML (Lan et al., 2024). In the period March 1958-December 1979, we use bias-adjusted values of the global growth rate based on measurements of atmospheric  $\text{CO}_2$  made by the Scripps Institution of

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Oceanography at the Mauna Loa Observatory, Hawaii (Keeling et al., 1976; full period of coverage 1758-2024). Bias adjustment of the Scripps data was performed in three sequential stages as follows:

- First, to correct for differences in the mean atmospheric concentration of CO<sub>2</sub> at Mauna Loa versus the globally averaged value, a constant of -0.231 ppm was added to all Scripps data to improve alignment of the “CO<sub>2</sub>[trend]” values from the Scripps data with the “CO<sub>2</sub>[trend]” values from the global NOAA data. The value of -0.231 ppm is the mean offset of “CO<sub>2</sub>[trend]” at Mauna Loa from the global “CO<sub>2</sub>[trend]” value during 1980-2000.
- Second, to correct for differences in the seasonality of atmospheric CO<sub>2</sub> concentrations at Mauna Loa versus globally, we shifted monthly anomalies between CO<sub>2</sub> concentration data and “trend” values backward in time by one month in the Scripps data. This specifically corrects for the fact that peaks/troughs in the climatology of “CO<sub>2</sub>[monthly observation] - CO<sub>2</sub>[trend]” at Mauna Loa occur 1 month earlier than peaks/troughs in the climatology of “CO<sub>2</sub>[monthly observation] - CO<sub>2</sub>[trend]” in the global data from NOAA. A one-month shift to the Scripps data was found to optimally align the climatologies of “CO<sub>2</sub>[monthly observation] - CO<sub>2</sub>[trend]” in the Scripps and global data.
- Third, to correct for the greater amplitude of seasonal anomalies at Mauna Loa from Scripps than the global data from NOAA, we apply a monthly multiplier that dampens the magnitude of monthly anomalies from “trend” values in the Scripps data. The monthly multiplier reduces values of “CO<sub>2</sub>[monthly observation] - CO<sub>2</sub>[trend]” in the Scripps data to more closely match values of “CO<sub>2</sub>[monthly observation] - CO<sub>2</sub>[trend]” in the NOAA global data.

For the period Jan 1750 to February 1958, we use bias-adjusted values of the global growth rate based on measurements of atmospheric CO<sub>2</sub> from air trapped in ice at Law Dome (Joos and Spahni, 2008; full period of coverage 1750-2004). Bias adjustments were made to improve alignment with the post-1980 time series of data from Scripps and NOAA, and were performed in two sequential stages as follows:

- First, a constant of 0.973 was added to all data from Law Dome to improve alignment with the Scripps data (which had already been bias-corrected as described above). The constant of 0.973 is the mean offset of CO<sub>2</sub> annual values (annual mean in the case of the Scripps data) in the period 1958-1979.
- Second, the climatology of “CO<sub>2</sub>[monthly observation] - CO<sub>2</sub>[trend]” from the period 1958-2000 was superimposed on the data from Law Dome (note that the 1958-2000 data includes both Scripps and NOAA data, combined as described above). To achieve this, a spline interpolation was fitted to downscale annual observations from CO<sub>2</sub> concentration from Law Dome to monthly values of “CO<sub>2</sub>[trend]” and the climatological seasonality of “CO<sub>2</sub>[monthly observation] - CO<sub>2</sub>[trend]” from 1958-2000) was then added to the interpolated values of “CO<sub>2</sub>[trend]”.

To derive S<sub>ocean</sub> from the model simulations, we subtracted the slope of a linear fit to the annual time series of the control simulation B from the annual time series of simulation A. Assuming that drift and bias are the same in simulations A and B, we thereby correct for any model drift. Further, this difference also removes the natural steady state flux (assumed to be 0 GtC yr<sup>-1</sup> globally without rivers), which is often a major source of biases.

Note, however, that Gürses et al. (2023) questioned the assumption of comparable bias and drift in simulations A and B as they compared two versions of FESOM-REcoM, and found a very similar air-sea CO<sub>2</sub> flux in

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simulation A despite a different bias as derived from simulation B. This approach works for all model set-ups, including IPSL, where simulation B was forced with [variable](#) historical [climate](#) changes ([looping over a 10-year forcing](#)). This approach assures that the interannual variability is not removed from IPSL simulation A. The absolute correction for bias and drift per model in the 1990s varied between  $<0.01 \text{ GtC yr}^{-1}$  and  $0.31 \text{ GtC yr}^{-1}$ , with five models having positive biases, four having negative biases and one model having essentially no bias (NorESM). The MPI model uses riverine input and therefore simulates outgassing in simulation B. By subtracting a linear fit of simulation B, also the ocean carbon sink of the MPI model follows the definition of  $\text{So}_{\text{CEAN}}$ . This correction [increases](#) the model mean ocean carbon sink by  $0.07 \text{ GtC yr}^{-1}$  in the 1990s. The ocean models cover 99% to 101% of the total ocean area, so that area-scaling is not necessary.

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### 848 S.3.3 GOBM evaluation

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849 The ocean CO<sub>2</sub> sink for all GOBMs and the ensemble mean falls within 90% confidence of the observed range, or 1.5 to 2.9 GtC yr<sup>-1</sup> for the 1990s (Ciais et al., 2013) before and after applying adjustments. The GOBMs and  $f\text{CO}_2$ -products have been further evaluated using the fugacity of sea surface CO<sub>2</sub> ( $f\text{CO}_2$ ) from the SOCAT [v2024](#) database (Bakker et al., 2016, [2024](#)). We focused this evaluation on the root mean squared error (RMSE) between observed and modelled  $f\text{CO}_2$  and on a measure of the amplitude of the interannual variability of the flux (modified after Rödenbeck et al., 2015). The RMSE is calculated from detrended, annually and regionally averaged time series of  $f\text{CO}_2$  calculated from GOBMs and  $f\text{CO}_2$ -products subsampled to SOCAT sampling points to measure the misfit between large-scale signals (Hauck et al., 2020). To this end, we apply the following steps: (i) subsample data points for where there are observations (GOBMs/ $f\text{CO}_2$ -products as well as SOCAT), (ii) average spatially, (iii) calculate annual mean, (iv) detrend both time-series (GOBMs/ $f\text{CO}_2$ -products as well as SOCAT), (v) calculate RMSE. We use a mask based on the minimum area coverage of the  $f\text{CO}_2$ -products. This ensures a fair comparison over equal areas. The amplitude of the  $\text{So}_{\text{CEAN}}$  interannual variability (A-IAV) is calculated as the temporal standard deviation of the detrended annual CO<sub>2</sub> flux time series after area-scaling (Rödenbeck et al., 2015, Hauck et al., 2020). These metrics are chosen because RMSE is the most direct measure of data-model mismatch and the A-IAV is a direct measure of the variability of  $\text{So}_{\text{CEAN}}$  on interannual timescales. We apply these metrics globally and by latitude bands. Results are shown in Figure S2 and discussed in Section 3.6.5.

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866 In addition to the interior ocean anthropogenic carbon accumulation (Section 3.6.5) and SOCAT  $f\text{CO}_2$ , we evaluate the models with process-based metrics that were previously related to ocean carbon uptake. These are the Atlantic Meridional Overturning Circulation (Goris et al., 2018, Terhaar et al., 2022, Terhaar et al., in review), the Southern Ocean sea surface salinity (Terhaar et al., 2021, 2022, [2024](#), Hauck et al., [2023b](#)), the Southern Ocean stratification index (Bourgeois et al., 2022) and the surface ocean Revelle factor (Terhaar et al., 2022, [2024](#)).

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873 We follow the methodology of previous studies wherever possible, particularly the RECCAP model evaluation chapter (Terhaar et al., [2024](#)). The Atlantic Meridional Overturning Circulation from the GOBMs is here defined as the maximum of the Atlantic meridional overturning streamfunction at 26°N. This is compared to data from

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the RAPID array at 26°N (Moat et al., 2024). An uncertainty of 0.9 Sv was reported in McCarthy et al. (2015). We use the years 2005–2022, which are all complete calendar years available from the RAPID data set, and report the temporal standard deviation over that period.

The Southern Ocean sea surface salinity is reported for the subpolar seasonally stratified biome (SPSS) and for the area covering both the SPSS and subtropical seasonally stratified (STSS) biomes. Biome definitions are taken from Fay and McKinley (2014, as provided for the RECCAP2 project). The sea surface salinity was first used as an emergent constraint for the Southern Ocean CO<sub>2</sub> uptake with Earth System Models (Terhaar et al. 2021, 2022) using the interfrontal salinity between the polar and subtropical fronts with dynamic fronts. As the GOBMs are forced with reanalysis data, the fronts do not vary as much as in the ESMs, and thus the use of fixed biomes is justified (Hauck et al., 2023b, Terhaar et al., 2024). We use the time period 2005–2022 for consistency with the AMOC metric. The observational sea surface salinity values are calculated from the EN4 data set (Good et al., 2013; using the objective analyses – Gouretski and Reseghetti (2010) XBT corrections and Gouretski and Cheng (2020) MBT corrections) with the aid of the Fay and McKinley (2014) mask.

The Southern Ocean stratification index is a simplified version of the metric used in Bourgeois et al. (2022). It is defined as the difference between in situ density at the surface and at 1000 m depth in the latitudinal band of 30°S to 55°S. Each model provider calculated this metric based on their native model mesh. We use again the period of 2005–2022 for consistency with the AMOC metric. The same metric was calculated from the EN4 data set mentioned above (Good et al., 2013).

Finally, the global surface ocean Revelle factor is reported. Monthly 1°x1° gridded fields were provided by the modelling groups, based on standard carbonate chemistry routines (e.g., mocsy, Orr & Epitalon, 2015; PyCO2SYS, Humphreys et al., 2022a,b). The observational metrics come from two sources, firstly the gridded GLODAP data set v2.2016 (Lauvset et al., 2016), which is a climatology centered around the year 2002. For comparison with GLODAP, the models were subsampled to GLODAP data coverage and to a comparable time window also centred around 2002 (1997–2007). Secondly, the OceanSODA v2024 data set (Gregor and Gruber, 2020, updated) was used, which has all input data available to calculate the surface ocean Revelle factor. OceanSODA covers a slightly smaller surface area (~96 % of GLODAP), but provides data until 2021. The period 2005–2021 was used due to data availability and the models were subsampled to the same spatial and temporal coverage.

For this release, only the comparison of the metrics between GOBMs and observational data sets is presented, whereas it is foreseen to translate this comparison into a quantitative benchmarking comparable to the iLAMB benchmarking for the DGVMs and the corresponding iOMB framework (Ogunro et al., 2018). In a next step, model weighting can be applied based on the benchmarking (e.g., Brunner et al., 2020).

### S3.4 /CO<sub>2</sub>-product trend benchmarking

In addition to the air-sea CO<sub>2</sub> flux estimates, /CO<sub>2</sub>-product providers reconstructed the sea surface /CO<sub>2</sub> of a set of 4 GOBM's, namely CESM-ETHZ, FESOM2.1REcoM, MRI-ESM2 and IPSL, that were submitted to the

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GCB2023 (Friedlingstein et al. 2023) following the approach of Hauck et al. (2023). A total of 6  $f\text{CO}_2$ -products conducted the benchmark test (VLIZ-SOMFFN, NIES-ML3, Jena-MLS, CSIR-ML6, OceanSODA-ETHZv2 and JMA-MLR). The GOBM's serve as known truth and are subsampled according to the real world observation tracks. The  $f\text{CO}_2$ -products then reconstruct the true model field, based on the subsampled information provided. We then compare trends for the period 2001-2021, i.e. the period where we see the divergence between  $f\text{CO}_2$ -products and models, removing the final year to avoid the tail effect. The trends of the individual  $f\text{CO}_2$ -products from the GCB24 were then plotted against the mean of the trend reconstruction bias (evaluated against the known truth GOBM trends) of the 4 GOBM. This is shown in Figure S3. The figure illustrates the tendency that  $f\text{CO}_2$ -products with negative biases in the  $f\text{CO}_2$  reconstruction show the strongest air-sea  $\text{CO}_2$  flux trends and vice versa for the  $f\text{CO}_2$  products with positive biases. Overall, the ensemble of 6  $f\text{CO}_2$  methods shows a tendency to underestimate the  $f\text{CO}_2$  trend from the GOBMs (with a mean bias across 6  $f\text{CO}_2$ -products and 4 model reconstructions of  $0.25 \mu\text{atm/decade}$ ) and thus an inferred tendency to overestimate the air-sea  $\text{CO}_2$  flux trend (mean across 6  $f\text{CO}_2$ -products of  $0.50 \pm 0.13 \text{ PgC yr}^{-1} \text{ decade}^{-1}$ ), however, due to compensating negative and positive  $f\text{CO}_2$  biases, the ensemble mean trend bias is smaller than suggested from previous studies focusing on one or two  $f\text{CO}_2$ -products only (see e.g. Gloege et al. 2021, Hauck et al. 2023). The inferred global trend of  $0.43 \pm 0.13 \text{ PgC yr}^{-1} \text{ decade}^{-1}$  that intercepts with the 0 bias line closely corresponds to a recent estimate by Mayot et al. 2024 of  $0.42 \pm 0.06 \text{ PgC yr}^{-1} \text{ decade}^{-1}$  (period 2000-2022) in the mean, although with a substantially larger uncertainty and different time period. The evidence basis, thus, remains low due to the small sample size of  $f\text{CO}_2$ -products ( $n=6$ ) and reconstructed GOBMs ( $n=4$ ), thus a more detailed analysis is required to better constrain  $f\text{CO}_2$ -product trends.

### S3.4 Uncertainty assessment for $\text{SOCEAN}$

We quantify the  $1-\sigma$  uncertainty around the mean ocean sink of anthropogenic  $\text{CO}_2$  by assessing random and systematic uncertainties for the GOBMs and  $f\text{CO}_2$ -products. The random uncertainties are taken from the ensemble standard deviation ( $0.3 \text{ GtC yr}^{-1}$  for GOBMs,  $0.3 \text{ GtC yr}^{-1}$  for  $f\text{CO}_2$ -products). We derive the GOBMs systematic uncertainty by the deviation of the DIC inventory change 1994-2007 from the Gruber et al. (2019) estimate ( $0.4 \text{ GtC yr}^{-1}$ ) and suggest these are related to physical transport (mixing, advection) into the ocean interior. For the  $f\text{CO}_2$ -products, we consider systematic uncertainties stemming from uncertainty in  $f\text{CO}_2$  observations ( $0.2 \text{ GtC yr}^{-1}$ , Takahashi et al., 2009; Wanninkhof et al., 2013), gas-transfer velocity ( $0.2 \text{ GtC yr}^{-1}$ , Ho et al., 2011; Wanninkhof et al., 2013; Roobaert et al., 2018), wind product ( $0.1 \text{ GtC yr}^{-1}$ , Fay et al., 2021), river flux adjustment ( $0.3 \text{ GtC yr}^{-1}$ , Regnier et al., 2022, formally  $2-\sigma$  uncertainty), and  $f\text{CO}_2$  mapping ( $0.2 \text{ GtC yr}^{-1}$ , Landschützer et al., 2014). Combining these uncertainties as their squared sums, we assign an uncertainty of  $\pm 0.5 \text{ GtC yr}^{-1}$  to the GOBMs ensemble mean and an uncertainty of  $\pm 0.6 \text{ GtC yr}^{-1}$  to the  $f\text{CO}_2$ -product ensemble mean, which is smaller than a recent estimate by Ford et al. (2024), who estimate an uncertainty of  $\pm 0.7 \text{ GtC yr}^{-1}$  based on propagating different sources of uncertainty in  $f\text{CO}_2$ -products. Here, the uncertainties are propagated as  $\sigma(\text{SOCEAN}) = (1/2^2 * 0.5^2 + 1/2^2 * 0.6^2)^{1/2} \text{ GtC yr}^{-1}$  and result in an  $\pm 0.4 \text{ GtC yr}^{-1}$  uncertainty around the best estimate of  $\text{SOCEAN}$ .

We examine the consistency between the variability of the GOBMs and the  $f\text{CO}_2$ -products to assess confidence in  $\text{SOCEAN}$ . The interannual variability of the ocean fluxes (quantified as A-IAV, the standard deviation after

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detrending, Figure S2) of the ~~eight~~  $f\text{CO}_2$ -products plus the UExP-FNN-U product (Watson et al., 2020; Ford et al., accepted) for 1990-2023, ranges from 0.08 to 0.37 GtC yr<sup>-1</sup> with the lower estimates by the three ensemble methods (NIES-ML3, CMEMS-LSCE-FFNN, OS-ETHZ-GRaCER). The inter-annual variability in the GOBMs ranges between 0.10 and 0.20 GtC yr<sup>-1</sup>, hence there is overlap with the A-IAV estimates of the  $f\text{CO}_2$ -products.

Individual estimates (both GOBMs and  $f\text{CO}_2$ products) generally produce a higher ocean CO<sub>2</sub> sink during strong El Niño events. There is emerging agreement between GOBMs and  $f\text{CO}_2$ -products on the patterns of decadal variability of SO<sub>CEAN</sub> with a global stagnation in the 1990s, an extra-tropical strengthening in the 2000s (McKinley et al., 2020, Hauck et al., 2020). More recently, a fast growth of the sink is simulated by both methods between 2001 and 2016, and a stagnation period since then. A stagnation or even decline of SO<sub>CEAN</sub> occurred during the triple La Niña years 2020-2023. The central estimates of the annual flux from the GOBMs and the  $f\text{CO}_2$ -products have a correlation  $r$  of 0.98 (1990-2023). The agreement between the models and the  $f\text{CO}_2$ products reflects some consistency in their representation of underlying variability since there is little overlap in their methodology or use of observations.

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## S.4 Methodology Land CO<sub>2</sub> sink

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### S.4.1 DGVM simulations

The DGVMs model runs were forced by either the merged monthly Climate Research Unit (CRU) and 6 hourly Japanese 55-year Reanalysis (JRA-55) data set or by the monthly CRU data set, both providing observation-based temperature, precipitation, and incoming surface radiation on a 0.5°x0.5° grid and updated to 2023 (Harris et al., 2014, 2020). The combination of CRU monthly data with 6 hourly forcing from JRA-55 (Kobayashi et al., 2015) is performed with methodology used in previous years (Viovy, 2016) adapted to the specifics of the JRA-55 data.

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Introduced in GCB2021 (Friedlingstein et al., 2022a), incoming short-wave radiation fields take into account aerosol impacts and the division of total radiation into direct and diffuse components as summarised below.

The diffuse fraction dataset offers 6-hourly distributions of the diffuse fraction of surface shortwave fluxes over the period 1901-2023. Radiative transfer calculations are based on monthly-averaged distributions of tropospheric and stratospheric aerosol optical depth, and 6-hourly distributions of cloud fraction. Methods follow those described in the Methods section of Mercado et al. (2009), but with updated input datasets.

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The time series of speciated tropospheric aerosol optical depth is taken from the historical and RCP8.5 simulations by the HadGEM2-ES climate model (Bellouin et al., 2011). To correct for biases in HadGEM2-ES, tropospheric aerosol optical depths are scaled over the whole period to match the global and monthly averages obtained over the period 2003-2020 by the CAMS Reanalysis of atmospheric composition (Inness et al., 2019), which assimilates satellite retrievals of aerosol optical depth.

The time series of stratospheric aerosol optical depth is taken from the by Sato et al. (1993) climatology, which has been updated to 2012. Years 2013-2020 are assumed to be background years so replicate the background year 2010. That assumption is supported by the Global Space-based Stratospheric Aerosol Climatology time series (1979-2016; Thomason et al., 2018). The time series of cloud fraction is obtained by scaling the 6-hourly distributions simulated in the Japanese Reanalysis (Kobayashi et al., 2015) to match the monthly-averaged cloud



cover in the CRU TS v4.06 dataset (Harris et al., 2020). Surface radiative fluxes account for aerosol-radiation interactions from both tropospheric and stratospheric aerosols, and for aerosol-cloud interactions from tropospheric aerosols, except mineral dust. Tropospheric aerosols are also assumed to exert interactions with clouds. The radiative effects of those aerosol-cloud interactions are assumed to scale with the radiative effects of aerosol-radiation interactions of tropospheric aerosols, using regional scaling factors derived from HadGEM2-ES. Diffuse fraction is assumed to be 1 in cloudy sky. Atmospheric constituents other than aerosols and clouds are set to a constant standard mid-latitude summer atmosphere, but their variations do not affect the diffuse fraction of surface shortwave fluxes.

In addition to the climate forcing, the DGVMs forcing also include the global atmospheric CO<sub>2</sub> time series, same as for the GOBMs and described in Section S.3.2 (Lan et al. (2023), the gridded land cover changes (see Supplement S.2.2), and the gridded nitrogen deposition and fertilisers (see Table S1 for specific models details). Four simulations were performed with each of the DGVMs. Simulation 0 (S0) is a control simulation which uses fixed pre-industrial (year 1700) atmospheric CO<sub>2</sub> concentrations, cycles early 20th century (1901-1920) climate and applies a time-invariant pre-industrial land cover distribution and pre-industrial wood harvest rates. Simulation 1 (S1) differs from S0 by applying historical changes in atmospheric CO<sub>2</sub> concentration and N inputs. Simulation 2 (S2) applies historical changes in atmospheric CO<sub>2</sub> concentration, N inputs, and climate, while applying time-invariant pre-industrial land cover distribution and pre-industrial wood harvest rates. Simulation 3 (S3) applies historical changes in atmospheric CO<sub>2</sub> concentration, N inputs, climate, and land cover distribution and wood harvest rates.

S2 is used to estimate the land sink component of the global carbon budget (S<sub>LAND</sub>). S3 is used to estimate the total land flux but is not used in the global carbon budget. We further separate S<sub>LAND</sub> into contributions from CO<sub>2</sub> (=S1-S0) and climate (=S2-S1+S0).

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#### S.4.2 DGVM evaluation

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We apply three criteria for minimum DGVMs realism by including only those DGVMs with (1) steady state after spin up, (2) global net land flux (S<sub>LAND</sub> – E<sub>LUC</sub>) that is an atmosphere-to-land carbon flux over the 1990s ranging between -0.3 and 2.3 GtC yr<sup>-1</sup>, within 90% confidence of constraints by global atmospheric and oceanic observations (Keeling and Manning, 2014; Wanninkhof et al., 2013), and (3) global E<sub>LUC</sub> that is a carbon source to the atmosphere over the 1990s, as already mentioned in Supplement S.2.2. All DGVMs meet these three criteria.

In addition, the DGVMs results are also evaluated using the International Land Model Benchmarking system (ILAMB; Collier et al., 2018). This evaluation is provided here to document, encourage and support model improvements through time. ILAMB variables cover key processes that are relevant for the quantification of S<sub>LAND</sub> and resulting aggregated outcomes (see Figure S4 for the results and for the list of observed databases). Results are shown in Figure S4 and briefly discussed in Section 3.7.5.

Deleted: . The selected variables are vegetation biomass, gross primary productivity, leaf area index, net ecosystem exchange, ecosystem respiration, evapotranspiration, soil carbon, runoff, and relationships between carbon cycle variables, precipitation (Adler et al., 2003) and temperature (Harris et al., 2014)

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The International LAnd Model Benchmarking (ILAMB) system (Collier et al. 2018; version 2.7.2 (2024): <https://github.com/rubisco-sfa/ILAMB/releases/tag/v2.7.2>) was used to compare the 21 models (20 DGVMs and CARDAMOM) to observational benchmarks for a number of different variables related to the land surface: gross primary productivity (GPP), leaf area index (LAI), ecosystem respiration, soil carbon, evapotranspiration,

runoff, burned areas, fire CO<sub>2</sub> emissions, and soil respiration), either for the entire global land surface or for the different RECCAP regions. Furthermore, relationships between selected pairs of variables can be visualised with ILAMB. Each row for each variable in Figs. S4 is clickable in the full website version [https://gws-access.jasmin.ac.uk/public/landsurf\\_rdg/pm McGuire/ILAMB\\_output/TRENDYv13\\_latest/](https://gws-access.jasmin.ac.uk/public/landsurf_rdg/pm McGuire/ILAMB_output/TRENDYv13_latest/) and gives access to geographic plots for such quantities as bias relative to observational benchmark, temporal RMSE from the observational benchmark, and difference in max month from the observational benchmark. The full website version also gives a spatial Taylor diagram for all the models, as well as time series comparisons of the regional mean time-series and the regional mean annual cycle. The Biomass variable was not included this year, due to a mismatch between the TRENDY *clveg* variable (above-ground and below-ground biomass, for all PFTs) and two of the previously-used observational benchmark datasets for biomass (Saatchi et al., 2011 and Thurner et al., 2014), which are both only for forests and for above-ground biomass. In the ILAMB setup for TRENDYv13, we have added three more variables (annual-averaged Burned Area, Fire Emissions, and Soil Respiration) and we have modified the Koven visualisation slightly for the Soil Carbon variable. All four of these changes have been put into a category of variables that we call ‘Ecosystem and Carbon Cycle Extended’. Two of the models (EDv3 and SDGVM) compute burned area either on a national level or without considering arid non-vegetated lands, as the model biases for burned area for these two models are rather high in the world’s deserts, compared to the GFED4.1S observational benchmark until the year 2016. However, in the case of SDGVM, the positive burned-area bias in the deserts is not apparent in the fire emissions variable. The Soil Respiration variable has been added only for those models that provided the *soilr* model output, which is calculated as the sum of heterotrophic respiration and root respiration. For the soil respiration variable, three observational benchmarks were selected (Tang et al. 2019, 2020, Raich et al. 2002 and Hashimoto et al. 2015) from the data sets contrasted by Hashimoto et al. (2023). The Koven analysis of the Soil Carbon turnover time is part of the standard setup in ILAMB version 2.7.2, but we put it into the Extended category largely since it seems to be missing proper application of an aridity mask for all of the models, unlike for the Observational Benchmark. We also added a model-fit curve to the Koven analysis, for better visualisation by allowing the comparison to the benchmark-fit curve. The TRENDYv13 version of the updated ILAMB version 2.7.2 GitHub code fork/branch is available at: <https://github.com/mcguirepatr/ILAMB/tree/master>

#### S.4.3 Uncertainty assessment for S<sub>LAND</sub>

For the uncertainty for S<sub>LAND</sub>, we use the standard deviation of the annual CO<sub>2</sub> sink across the DGVMs, averaging to about ± 0.6 GtC yr<sup>-1</sup> for the period 1959 to 2021. We attach a medium confidence level to the annual land CO<sub>2</sub> sink and its uncertainty because the estimates from the residual budget and averaged DGVMs match well within their respective uncertainties (Table 5).

### S.5 Methodology Atmospheric Inversions

#### S.5.1 Inversion System Simulations

Fourteen atmospheric inversions (details of each in Table S4) were used to infer the spatio-temporal distribution of the CO<sub>2</sub> flux exchanged between the atmosphere and the land or oceans. These inversions are based on

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1142 Bayesian inversion principles with prior information on fluxes and their uncertainties. They use very similar sets  
 1143 of surface measurements of CO<sub>2</sub> time series (or subsets thereof) from various flask and in situ networks. Six  
 1144 inversion systems used satellite xCO<sub>2</sub> retrievals from GOSAT and/or OCO-2, of which two systems used a  
 1145 combination of satellite and surface observations.  
 1146 Each inversion system uses different methodologies and input data but is rooted in Bayesian inversion  
 1147 principles. These differences mainly concern the selection of atmospheric CO<sub>2</sub> data and prior fluxes, as well as  
 1148 the spatial resolution, assumed correlation structures, and mathematical approach of the models. Each system  
 1149 uses a different transport model, which was demonstrated to be a driving factor behind differences in  
 1150 atmospheric inversion-based flux estimates, and specifically their distribution across latitudinal bands (Gaubert  
 1151 et al., 2019; Schuh et al., 2019).  
 1152 Most of the fourteen inversion systems prescribe similar global fossil fuel emissions for E<sub>FOS</sub>; specifically, the  
 1153 GCP's Gridded Fossil Emissions Dataset version 2024.0 (GCP-GridFEDv2024.0; Jones et al., 2024), which is  
 1154 an update through 2023 of the first version of GCP-GridFED presented by Jones et al. (2021b) (Table S4). All  
 1155 GCP-GridFED versions scale gridded estimates of CO<sub>2</sub> emissions from EDGARv4.3.2 (Janssens-Maenhout et  
 1156 al., 2019) within national territories to match national emissions estimates provided by the GCP for the years  
 1157 1959-2023, which are compiled following the methodology described in Supplement S.1. GCP-  
 1158 GridFEDv2024.0 adopts the seasonality of emissions (the monthly distribution of annual emissions) from the  
 1159 Carbon Monitor (Liu et al., 2020a,b; Dou et al., 2022) for Brazil, China, all EU27 countries, the United  
 1160 Kingdom, the USA and shipping and aviation bunker emissions. The seasonality present in Carbon Monitor is  
 1161 used directly for years 2019-2023, while for years 1959-2018 the average seasonality of 2019, and 2021 and  
 1162 2022 are applied (avoiding the year 2020 during which emissions were most impacted by the COVID-19  
 1163 pandemic). For all other countries, seasonality of emissions is taken from EDGAR (Janssens-Maenhout et al.,  
 1164 2019; Jones et al., 2023), with small annual correction to the seasonality present in 2010 based on heating or  
 1165 cooling degree days to account for the effects of inter-annual climate variability on the seasonality of emissions  
 1166 (Jones et al., 2021b).  
 1167 Small remaining differences between regridding of the GridFED inputs, or the use of different fossil fuel  
 1168 emission priors are corrected for by scaling the resulting inverse fluxes to GridFEDv2024.0. The consistent use  
 1169 of E<sub>FOS</sub> ensures a close alignment with the estimate of E<sub>FOS</sub> used in this budget assessment, enhancing the  
 1170 comparability of the inversion-based estimate with the flux estimates deriving from DGVMs, GOBMs and  
 1171 /CO<sub>2</sub>-based methods. The fossil fuel adjustment (including emissions from cement production and cement  
 1172 carbonation CO<sub>2</sub> sink) ensures that the estimated uptake of atmospheric CO<sub>2</sub> by the land and oceans was fully  
 1173 consistent within the inversion ensemble.  
 1174 The land and ocean CO<sub>2</sub> fluxes from atmospheric inversions contain anthropogenic perturbation and natural pre-  
 1175 industrial CO<sub>2</sub> fluxes. On annual time scales, natural pre-industrial fluxes are primarily land CO<sub>2</sub> sinks and  
 1176 ocean CO<sub>2</sub> sources corresponding to carbon taken up on land, transported by rivers from land to ocean, and  
 1177 outgassed by the ocean. These pre-industrial land CO<sub>2</sub> sinks are thus compensated over the globe by ocean CO<sub>2</sub>  
 1178 sources corresponding to the outgassing of riverine carbon inputs to the ocean, using the exact same numbers  
 1179 and distribution as described for the oceans in Section 2.5. To facilitate the comparison, we adjusted the inverse  
 1180 estimates of the land and ocean fluxes per latitude band with these numbers to produce historical perturbation  
 1181 CO<sub>2</sub> fluxes from inversions.

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## 1191 S.5.2 Inversion System Evaluation

1192 All participating atmospheric inversions are checked for consistency with the annual global growth rate, as both  
 1193 are derived from the global surface network of atmospheric CO<sub>2</sub> observations. In this exercise, we use the  
 1194 conversion factor of 2.086 GtC/ppm to convert the inverted carbon fluxes to mole fractions, as suggested by  
 1195 Prather (2012). This number is specifically suited for the comparison to surface observations that do not respond  
 1196 uniformly, nor immediately, to each year's summed sources and sinks. This factor is therefore slightly smaller  
 1197 than the GCB conversion factor in Table 1 (2.142 GtC/ppm, Ballantyne et al., 2012). Overall, the inversions  
 1198 agree with the growth rate with biases between 0.0002-0.065 ppm yr<sup>-1</sup> (0.0004-0.13 GtCyr<sup>-1</sup>) for the period  
 1199 2015-2023.

1200 The atmospheric inversions are also evaluated using vertical profiles of atmospheric CO<sub>2</sub> concentrations (Figure  
 1201 S5). More than 30 aircraft programs over the globe, either regular programs or repeated surveys over at least 9  
 1202 months (except on the SH), have been used in order to draw a robust picture of the system performance (with  
 1203 space-time data coverage irregular and denser in the 0-45°N latitude band; Table S8 and lower panel in Figure  
 1204 S4). The fourteen systems are compared to these independent aircraft CO<sub>2</sub> observations between 2 and 7 km  
 1205 above sea level between 2001 and 2023. Results are shown in Figure S5, where the inversions generally match  
 1206 the atmospheric mole fractions to within 0.7 ppm at all latitudes.

1207

## 1208 S.6 Processes not included in the global carbon budget

### 1209 S.6.1 Contribution of anthropogenic CO and CH<sub>4</sub> to the global carbon budget

1210 Equation (1) includes only partly the net input of CO<sub>2</sub> to the atmosphere from the chemical oxidation of reactive  
 1211 carbon-containing gases from sources other than the combustion of fossil fuels, such as: (1) cement process  
 1212 emissions, since these do not come from combustion of fossil fuels, (2) the oxidation of fossil fuels, (3) the  
 1213 assumption of immediate oxidation of vented methane in oil production. However, it omits any other  
 1214 anthropogenic carbon-containing gases that are eventually oxidised in the atmosphere, forming a diffuse source  
 1215 of CO<sub>2</sub>, such as anthropogenic emissions of CO and CH<sub>4</sub>. An attempt is made in this section to estimate their  
 1216 magnitude and identify the sources of uncertainty. Anthropogenic CO emissions are from incomplete fossil fuel  
 1217 and biofuel burning and deforestation fires. The main anthropogenic emissions of fossil CH<sub>4</sub> that matter for the  
 1218 global (anthropogenic) carbon budget are the fugitive emissions of coal, oil and gas sectors (see below). These  
 1219 emissions of CO and CH<sub>4</sub> contribute a net addition of fossil carbon to the atmosphere.

1220 In our estimate of E<sub>FOS</sub> we assumed (Section 2.1.1) that all the fuel burned is emitted as CO<sub>2</sub>, thus CO  
 1221 anthropogenic emissions associated with incomplete fossil fuel combustion and its atmospheric oxidation into  
 1222 CO<sub>2</sub> within a few months are already counted implicitly in E<sub>FOS</sub> and should not be counted twice (same for ELUC  
 1223 and anthropogenic CO emissions by deforestation fires). The diffuse atmospheric source of CO<sub>2</sub> deriving from  
 1224 anthropogenic emissions of fossil CH<sub>4</sub> is not included in E<sub>FOS</sub>. In reality, the diffuse source of CO<sub>2</sub> from CH<sub>4</sub>  
 1225 oxidation contributes to the annual CO<sub>2</sub> growth. Emissions of fossil CH<sub>4</sub> represent 30% of total anthropogenic  
 1226 CH<sub>4</sub> emissions (Saunio et al. 2020; their top-down estimate is used because it is consistent with the observed  
 1227 CH<sub>4</sub> growth rate), that is 0.083 GtC yr<sup>-1</sup> for the decade 2008-2017. Assuming steady state, an amount equal to

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Deleted: , except for MIROC4-ACTM in the Northern Hemisphere in the 2015-2022 period. Based on this larger bias with also larger standard deviations, plus the larger bias for the growth rate, the results for MIROC4-ACTM are not included in the statistics of the inversion ensemble

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1243 this fossil CH<sub>4</sub> emission is all converted to CO<sub>2</sub> by OH oxidation, and thus explain 0.083 GtC yr<sup>-1</sup> of the global  
1244 CO<sub>2</sub> growth rate with an uncertainty range of 0.061 to 0.098 GtC yr<sup>-1</sup> taken from the min-max of top-down  
1245 estimates in Saunio et al. (2020). If this min-max range is assumed to be 2  $\sigma$  because Saunio et al. (2020) did  
1246 not account for the internal uncertainty of their min and max top-down estimates, it translates into a 1- $\sigma$   
1247 uncertainty of 0.019 GtC yr<sup>-1</sup>.

1248 Other anthropogenic changes in the sources of CO and CH<sub>4</sub> from wildfires, vegetation biomass, wetlands,  
1249 ruminants, or permafrost changes are similarly assumed to have a small effect on the CO<sub>2</sub> growth rate. The CH<sub>4</sub>  
1250 and CO emissions and sinks are published and analysed separately in the Global Methane Budget and Global  
1251 Carbon Monoxide Budget publications, which follow a similar approach to that presented here (Saunio et al.,  
1252 2020; Zheng et al., 2019).

1253

## 1254 S.6.2 Contribution of other carbonates to CO<sub>2</sub> emissions

1255 Although we do account for cement carbonation (a carbon sink), the contribution of emissions of fossil  
1256 carbonates (carbon sources) other than cement production is not systematically included in estimates of E<sub>FOS</sub>,  
1257 except for Annex I countries and lime production in China (Andrew and Peters, 2021). The missing processes  
1258 include CO<sub>2</sub> emissions associated with the calcination of lime and limestone outside of cement production.  
1259 Carbonates are also used in various industries, including in iron and steel manufacture and in agriculture. They  
1260 are found naturally in some coals. CO<sub>2</sub> emissions from fossil carbonates other than cement not included in our  
1261 dataset are estimated to amount to about 0.3% of E<sub>FOS</sub> (estimated based on Crippa et al., 2019).

1262

## 1263 S.6.3 Anthropogenic carbon fluxes in the land-to-ocean aquatic continuum

1264 The approach used to determine the global carbon budget refers to the mean, variations, and trends in the  
1265 perturbation of CO<sub>2</sub> in the atmosphere, referenced to the pre-industrial era. Carbon is continuously displaced  
1266 from the land to the ocean through the land-ocean aquatic continuum (LOAC) comprising freshwaters, estuaries,  
1267 and coastal areas (Bauer et al., 2013; Regnier et al., 2013). A substantial fraction of this lateral carbon flux is  
1268 entirely ‘natural’ and is thus a steady state component of the pre-industrial carbon cycle. We account for this  
1269 pre-industrial flux where appropriate in our study (see Supplement S.3). However, changes in environmental  
1270 conditions and land-use change have caused an increase in the lateral transport of carbon into the LOAC – a  
1271 perturbation that is relevant for the global carbon budget presented here.

1272 The results of the analysis of Regnier et al. (2013) can be summarised in two points of relevance for the  
1273 anthropogenic CO<sub>2</sub> budget. First, the anthropogenic perturbation of the LOAC has increased the organic carbon  
1274 export from terrestrial ecosystems to the hydrosphere by as much as  $1.0 \pm 0.5$  GtC yr<sup>-1</sup> since pre-industrial  
1275 times, mainly owing to enhanced carbon export from soils. Second, this exported anthropogenic carbon is partly  
1276 respired through the LOAC, partly sequestered in sediments along the LOAC and to a lesser extent, transferred  
1277 to the open ocean where it may accumulate or be outgassed. The increase in storage of land-derived organic  
1278 carbon in the LOAC carbon reservoirs (burial) and in the open ocean combined is estimated by Regnier et al.  
1279 (2013) at  $0.65 \pm 0.35$  GtC yr<sup>-1</sup>. The inclusion of LOAC related anthropogenic CO<sub>2</sub> fluxes should affect estimates  
1280 of S<sub>LAND</sub> and S<sub>OCEAN</sub> in Eq. (1) but does not affect the other terms. Representation of the anthropogenic

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1281 perturbation of LOAC CO<sub>2</sub> fluxes is however not included in the GOBMs and DGVMs used in our global  
1282 carbon budget analysis presented here.

1283

1284 **S.6.4 Loss of additional land sink capacity**

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1285 Historical land-cover change was dominated by transitions from vegetation types that can provide a large carbon  
1286 sink per area unit (typically, forests) to others less efficient in removing CO<sub>2</sub> from the atmosphere (typically,  
1287 croplands). The resultant decrease in land sink, called the 'loss of additional sink capacity', can be calculated as  
1288 the difference between the actual land sink under changing land-cover and the counterfactual land sink under  
1289 pre-industrial land-cover. This term is not accounted for in our global carbon budget estimate. Here, we provide  
1290 a quantitative estimate of this term to be used in the discussion. Seven of the DGVMs used in Friedlingstein et  
1291 al. (2019) performed additional simulations with and without land-use change under cycled pre-industrial  
1292 environmental conditions. The resulting loss of additional sink capacity amounts to  $0.9 \pm 0.3$  GtC yr<sup>-1</sup> on  
1293 average over 2009-2018 and  $42 \pm 16$  GtC accumulated between 1850 and 2018 (Obermeier et al., 2021).  
1294 OSCAR, emulating the behaviour of 11 DGVMs finds values of the loss of additional sink capacity of  $0.7 \pm 0.6$   
1295 GtC yr<sup>-1</sup> and  $31 \pm 23$  GtC for the same time period (Gasser et al., 2020). Since the DGVM-based ELUC  
1296 estimates are only used to quantify the uncertainty around the bookkeeping models' ELUC, we do not add the  
1297 loss of additional sink capacity to the bookkeeping estimate.

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## Supplementary Tables

**Table S1.** Comparison of the processes included in the bookkeeping method and DGVMs in their estimates of ELUC and SLAND. See Table 4 for model references. All models include deforestation and forest regrowth after abandonment of agriculture (or from afforestation activities on agricultural land). Processes relevant for ELUC are only described for the DGVMs used with land-cover change in this study.

	Bookkeeping Models				DGVMs																								
	H&C2023	BLUE	OSCAR	LUCE	CABLE-POP	CLAS-SIC	CLM6.0	DLEM	EDv3	ELM	IBIS	IMAPLE	ISAM	ISBA-CTRIPE	JSBA-CH	JULES-ES	LPJ-GUESS	LPJml	LPJws	LPX-Bern	OCNv2	ORCHIDEEv3	SDGM	VISIT					
<b>Processes relevant for ELUC</b>																													
Wood harvest and forest degradation (a)	yes	yes	yes	yes	yes	yes	no	yes	yes	yes	yes	no	yes	no	yes	no	yes	no	yes	no (d)	yes	yes	yes	yes	yes	yes	yes	yes	yes
Shifting cultivation / Subgrid scale transitions	yes (b)	yes	yes	yes	yes	yes	no	yes	no	yes	yes	yes	no	no	yes	no	yes	no	yes	no (d)	no	no	yes	yes	yes	yes	yes	yes	yes
Cropland harvest (removed, R, or added to litter, L)	yes (R) (j)	yes (R) (j)	yes (R)	yes (R) (j)	yes (R) (j)	yes (R) (L)	yes (R+L)	yes	yes (R+L)	yes (L)	yes (R)	yes (L)	yes	yes (R)	yes (R+L)	yes (R)	yes (R)	yes (R+L)	yes (L)	yes (R)	yes (R+L)	yes (R)	yes (R)	yes (R)	yes (R)	yes (R)	yes (R)	yes (R)	yes (R)
Peat fires	yes (k)	yes (k)	yes (k)	yes (k)	yes (k)	no	no	yes	no	no	no	no	no	yes	no	no	no	no	no	no	no	no	no	no	no	no	no	no	no
fire as a management tool	yes (j)	yes (j)	yes (h)	yes (j)	yes (j)	no	no	no	no	no	no	no	no	no	no	no	no	no	no	no	no	no	no	no	no	no	no	no	no
N fertilisation	yes (j)	yes (j)	yes (h)	yes (j)	yes (j)	no	no	yes	yes	no	no	yes	no	yes	no	yes (i)	yes	yes	no	yes	yes	yes	yes	no	no	no	no	no	no
tillage	yes (j)	yes (j)	yes (h)	yes (j)	yes (j)	no	yes (g)	yes	yes	no	no	no	no	no	no	no	yes	yes	no	no	no	yes (g)	no	no	no	no	no	no	no
irrigation	yes (j)	yes (j)	yes (h)	yes (j)	yes (j)	no	no	yes	yes	no	no	no	yes	no	no	no	yes	yes	no	no	no	no	no	no	no	no	no	no	no
wetland drainage	yes (j)	yes (j)	yes (h)	yes (j)	yes (j)	no	no	no	no	no	no	no	yes	no	no	no	no	no	no	no	no	no	no	no	no	no	no	no	no
erosion	yes (j)	yes (j)	yes (h)	yes (j)	yes (j)	no	no	no	yes	no	no	yes	no	no	no	no	no	no	no	no	no	no	no	no	yes	yes	yes	yes	yes





**Table S2. Comparison of the processes and model set up for the Global Ocean Biogeochemistry Models for their estimates of SOCEAN. See Table 4 for model references.**

	NEMO-PlankTOM12	NEMO4.2-PISCES (IPSL)	MICOM-HAMOC (NorESM1-OCv1.2)	MPIOM-HAMOC6	FESOM-2.1-REcoM3	NEMO3.6-PISCESv2-gas (CNRM)	MOM6-COBALT (Princeton)	CESM-ETHZ	MRI-ESM2-3	ACCESS (CSIRO)
<b>Model specifics</b>										
Physical ocean model	NEMOv3.6-ORCA2	NEMOv4.2-eORCA1L75	MICOM (NorESM1-OCv1.2)	MPIOM	FESOM-2.1	NEMOv3.6-GELATov6-eORCA1L75	MOM6-SIS2	CESMv1.3 (ocean model based on POP2)	MRI.COMv5	MOM5
Biogeochemistry model	PlankTOM12	PISCESv2	HAMOC (NorESM1-OCv1.2)	HAMOC6	REcoM-3	PISCESv2-gas	COBALTv2	BEC (modified & extended)	NPZD+Fe	WOMBAT
Horizontal resolution	2° lon, 0.3 to 1.5° lat	1° lon, 0.3 to 1° lat	1° lon, 0.17 to 0.25 lat	1.5°	unstructured mesh, 20-120 km resolution (CORE mesh)	1° lon, 0.3 to 1° lat	0.5° lon, 0.25 to 0.5° lat	1.125° lon, 0.53° to 0.27° lat	1° lon, 0.3 to 0.5° lat	1°x1° with enhanced latitudinal resolution in the tropics and lat Southern Ocean
Vertical resolution	31 levels	75 levels, 1m at the surface	51 isopycnal layers + 2 layers representing a bulk mixed layer	40 levels	46 levels, 10 m spacing in the top 100 m	75 levels, 1m at surface	75 levels hybrid coordinates, 2m at surface	60 levels	60 levels with 1-level bottom boundary layer	50 levels, 20 m in the top 200m
Total ocean area on native grid (km2)	3.6080E+08	3.6360E+08	3.6006E+08	3.6598E+08	3.6435E+08	3.6270E+14	3.6111E+08	3.5926E+08	3.6094E+08	3.613E+08
Gas-exchange parameterization	Wanninkhof et al (1992)	Orr et al., 2017	Orr et al., 2017 but with a=0.337	Orr et al., 2017	Orr et al., 2017	Orr et al., 2017	Wanninkhof et al., 2014	Wanninkhof (1992, coefficient a scaled down to 0.31)	Orr et al., 2017	Wanninkhof et al (1992)
CO2 chemistry routines	OCMIP2 (Orr et al. 2017)	mocsy	Following Dickson et al., 2007	Ilyina et al. (2013) adapted to comply with OMIP protocol (Orr et al., 2017)	mocsy	mocsy	mocsy	OCMIP2 (Orr et al. 2017)	mocsy	OCMIP2 (Orr et al. 2017)
River input (PgC/yr) (organic/inorganic DIC)	0.723 / -	0.9167 (0.2577 / 0.659)	0	0.77 / -	0 / 0	0.611 / -	~0.07 / ~0.15	0.33 / -	0 / 0	0/0
Net flux to sediment (PgC/yr) (organic/other)	0.723 / -	0.3969 (0.0855 / 0.3114)	around 0.54 / -	0.71 / -	0 / 0	around 0.656 / -	~0.11 / ~0.07 (CaCO3)	0.21 / -	0 / 0	0/0
<b>SPIN-UP procedure</b>										
Initialisation of carbon chemistry	GLODAPv2 (preindustrial DIC)	GLODAPv2 (preindustrial DIC)	GLODAPv1 (preindustrial DIC)	initialization from previous simulation	GLODAPv2 (preindustrial DIC)	GLODAPv2	GLODAPv2 (Alkalinity, DIC). DIC corrected to 1959 level (simulation A and C) and to pre-industrial level (simulation B and C)	GLODAPv2 (preindustrial DIC)	GLODAPv2 (preindustrial DIC)	GLODAPv1 preindustrial DIC



	Jena-MLS	VLIZ-SOMFFN	CMEMS-LSCE-FFNN	UEXP-FNN-U (previously Watson et al.)	NIES-ML3	JMA-MLR	OceanSODA-ETHZv2	LDEO HPD	CSIR-ML6
Method	Spatio-temporal interpolation (version oc_v2023). Spatio-temporal field of ocean-internal carbon sources/sinks is fit to the SOCATv2022 pCO2 data. Includes a multi-linear regression against environmental drivers to bridge data gaps,	A feed-forward neural network (FFN) determines non-linear relationship between SOCAT pCO2 measurements and environmental predictor data for 16 biogeochemical provinces (defined through a self-organizing map, SOM) and is used to fill the existing data gaps.	An ensemble of neural network models trained on 100 subsampled datasets from SOCAT and environmental predictors. The models are used to reconstruct sea surface fugacity of CO2 and convert to air-sea CO2 fluxes	A self organising map feed forward neural network (SOM-FNN) implementation using SOCATv2024 fCO2 database, corrected to the subskin temperature (ESA CCI v3 bias corrected to surface drifter data following recommendations in Dong et al. 2022) of the ocean as measured by satellites (Goddijn-Murphy et al. 2015). Flux calculation corrected for the cool and salty surface skin. Monthly skin temperature	The ensemble of a random forest, a gradient boost machine, and a feed forward neural network trained on SOCAT fCO2 and environmental predictor variables. The interannual trend of fCO2 was estimated first by the decadal trend of atmospheric CO2 and then corrected by a so-called leave-one-year-out validation method. The trend was used to normalize fCO2 to the mid year of 1982-2023 for model training. The monthly fCO2	Fields of total alkalinity (TA) were estimated by using a multiple linear regressions (MLR) method based on GLODAPv2. 2023 and satellite observation data. SOCATv2024 fCO2 data were converted to dissolved inorganic carbon (DIC) with the TA. Fields of DIC were estimated by using a MLR method based on the DIC and satellite observation data	OceanSODA-ETHZv2 is a two-phase machine learning approach. In phase 1, we estimate the fCO2 8-day seasonal cycle climatology with a Gradient Boosted Decision Tree which is used as a predictor in the next phase. In phase 2, we predict the non-thermal component of fCO2 at a 8-day by 0.25° by 0.25° resolution with a two-layer fully-connected neural network using 35 ensemble members. The atmospheric CO2	Based on fCO2-misfit between observed fCO2 and 10 Global Carbon BudgetGOBMs. The eXtreme Gradient Boosting method links this misfit to environmental observations to reconstruct the model misfit across all space and time., which is then added back to the model-based fCO2 estimate. The final reconstruction of surface fCO2 is the average across the 10 reconstructions. A climatology of the misfits calculated for the years	An ensemble average of six machine-learning models, where each model is constructed with a two-step clustering-regression approach to determine a non-linear relationship between SOCAT fCO2 measurements and environmental proxy variables, and it used to fill the existing data gaps. The clustering step consists of two methods: the Mini-batch K-means clustering and the extended Fay and McKinley

				calculated from ESA CCI <a href="#">v3</a> ( <a href="#">Embury et al. 2024</a> ) with the cool skin difference calculated using NOAA COARE 3.5. <a href="#">Flux calculations completed using FluxEngine</a> ( <a href="#">Shutler et al., 2016</a> ; <a href="#">Holding et al., 2019</a> ).	maps were reconstructed using <a href="#">model prediction and</a> the trend.		<a href="#">and non-thermal component are added back to the result.</a>	2000-2023 is used as an offset for years prior to 1982 when no/limited environmental observations are available to train the ML algorithm.	(2014) biomes. The regression step consists of three methods: Gradient Boosting Machine, Support Vector Regression and Feed-forward Neural Network.
Gas-exchange parameterization	Wanninkhof 1992. Transfer coefficient k scaled to match a global mean transfer rate of 16.5 cm/hr by (Naegler, 2009)	Wanninkhof 1992. Transfer coefficient k scaled to match a global mean transfer rate of 16.5 cm/hr	Wanninkhof 2014. Transfer coefficient k scaled to match a global mean transfer rate of 16.5 cm/hr (Naegler, 2009)	Nightingale et al. 2000.	Wanninkhof, 2014. Transfer coefficient k scaled to match a global mean transfer rate of 16.5 cm/hr in <a href="#">1990-2019 (Fay et al., 2021)</a>	Wanninkhof, 2014. Transfer coefficient k scaled to match a global mean transfer rate of 16.5 cm/hr (Naegler, 2009)	Wanninkhof 1992. averaged and scaled for three reanalysis wind data, to a global mean 16.5 cm/hr (after Naegler 2009; Fay & Gregor et al. 2021)	Wanninkhof 1992. parameterization. Transfer coefficient k scaled to match a global mean transfer rate of 16.5 cm/hr (Naegler, 2009)	Wanninkhof 1992. averaged and scaled for three reanalysis wind data, to a global mean 16.5 cm/hr (after Naegler 2009; Fay & Gregor et al. 2021)
Wind product	JMA55-do reanalysis	ERA 5	ERA5	CCMP3.1	ERA5	JRA3Q	ERA5	ERA5	ERA 5
Spatial resolution	2.5 degrees longitude x 2 degrees latitude	1x1 degree	0.25x0.25 degree regridded to 1x1 degree	1x1 degree	<a href="#">Regrid 0.25x0.25 degree monthly data to 1x1</a>	1x1 degree	<a href="#">0.25x0.25 degree regridded to 1x1 degree</a>	1x1 degree	<a href="#">1x1 degree</a>

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Temporal resolution	daily	monthly	monthly	monthly	monthly	monthly	<a href="#">8-daily regridded to monthly</a>	monthly	monthly
Atmospheric CO2	Spatially and temporally varying field based on atmospheric CO2 data from 169 stations (Jena CarboScope atmospheric inversion sEXTALL_v2021)	Spatially varying 1x1 degree atmospheric pCO2_wet calculated from the NOAA <a href="#">ESRL</a> marine boundary layer xCO2 and NCEP sea level pressure with the moisture correction by Dickson et al 2007.	Spatially and monthly varying fields of atmospheric pCO2 computed from CO2 mole fraction (CO2 atmospheric inversion from the Copernicus Atmosphere Monitoring Service), and atmospheric dry-air pressure which is derived from monthly surface pressure (ERA5) and water vapour pressure fitted by Weiss and Price 1980.	Atmospheric fCO2 (wet) calculated from NOAA marine boundary layer XCO2(atm) and ERA5 sea level pressure, with pH2O calculated from Cooper et al. (1998). <a href="#">2023</a> XCO2 marine boundary values were not available at submission so we used preliminary values, estimated from <a href="#">2022</a> values and increase at Mauna Loa.	NOAA Greenhouse Gas Marine Boundary Layer Reference. <a href="https://gml.noaa.gov/ccgg/mb/mb.html">https://gml.noaa.gov/ccgg/mb/mb.html</a>	Atmospheric xCO2 fields of JMA-GSAM inversion model (Maki et al. 2010; Nakamura et al. 2015) were converted to pCO2 by using <a href="#">JRA30</a> sea level pressure. <a href="#">2023</a> xCO2 fields were not available at this stage, and we used Cape Grim and Mauna Loa xCO2 increments from <a href="#">2022</a> to <a href="#">2023</a> for the southern and northern hemispheres, respectively.	NOAA's marine boundary layer product for xCO2 is linearly interpolated onto a <a href="#">0.25°x0.25°</a> grid and resampled from weekly to <a href="#">8-daily</a> . xCO2 is multiplied by ERA5 mean sea level pressure, where the latter corrected for water vapour pressure using seasonal climatologic xCO2 values for a subset of common years (1979- <del>1989</del> ) yields a mean seasonality difference which is then applied to the Mauna Loa time series. Monthly 1x1 degree xCO2 is multiplied by	NOAA's marine boundary layer (MBL) surface xCO2 product is linearly interpolated to a 1x1 degree monthly grid for years 1979-2023. Prior to 1979, calculating an offset between the MBL and Mauna Loa, where the latter corrected for water vapour pressure using seasonal climatologic xCO2 values for a subset of common years (1979- <del>1989</del> ) yields a mean seasonality difference which is then applied to the Mauna Loa time series. Monthly 1x1 degree xCO2 is multiplied by	NOAA's marine boundary layer product for xCO2 is linearly interpolated onto a 1x1 degree monthly grid for years 1979-2023. Prior to 1979, calculating an offset between the MBL and Mauna Loa, where the latter corrected for water vapour pressure using seasonal climatologic xCO2 values for a subset of common years (1979- <del>1989</del> ) yields a mean seasonality difference which is then applied to the Mauna Loa time series. Monthly 1x1 degree xCO2 is multiplied by

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								ERA5 mean sea level pressure, with the correction for water vapour pressure using Dickson et al. 2007, using ERA5 SST and EN4 SSS. Finally converted to fCO2 using ERA5 SST and SLP.	
Total ocean area on native grid (km2)	3.63E+08	3.63E+08	3.50E+08	3.61E+08	3.19E+08	3.55E+08	3.586 E+8	3.63E+08	
method to extend product to full global ocean coverage		Arctic and marginal seas added following Landschützer et al. (2020). No coastal cut.			Fay & Gregor et al. 2021	Coverage of the global ice free ocean (ice frac < 0.9)	Based on method in Fay & Gregor et al. 2021. Gaps were filled with monthly climatology (Landschützer et al. 2020) scaled for interannual variability based on the temporal evolution of this product for all years.	Method has near full coverage	

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**Table S4.** Comparison of the inversion set up and input fields for the atmospheric inversions. Atmospheric inversions see the full CO2 fluxes, including the anthropogenic and pre-industrial fluxes. [Hence](#), they need to be adjusted for the pre-industrial flux of CO2 from the land to the ocean that is part of the natural carbon cycle before they can be compared with SOCEAN and SLAND from process models. See Table 4 for references.

Name	Jena CarboScope	Copernicus Atmosphere Monitoring Service (CAMS)	Carbon- Tracker Europe (CTE)	NISMON- CO2	CT-NOAA	CMS-Flux	Copernicus Atmosphere Monitoring Service (CAMS)	GONGGA	COLA	GCASv2	UoE	IAPCAS	MIROC- ACTM	NTFVAR
Version number	<a href="#">v76nbtEXToc v2024E</a>	<a href="#">v23r1</a>	<a href="#">v2024</a>	<a href="#">v2024_1</a>	CT2022 + CT- NRT <a href="#">v2024-1</a>	<a href="#">v2024</a>	<a href="#">FT24r1</a>	v2023	<a href="#">v2024</a>	<a href="#">v2024</a>	<a href="#">v2024</a>	<a href="#">v2024</a>	<a href="#">v2024</a>	<a href="#">v2024</a>
Flags														
Observations														
Atmospheric observations <a href="#">(a, b)</a>	Flasks and hourly from various institutions (outliers removed by 2σ criterion)	Hourly resolution (well-mixed conditions) obspack GLOBALVIE Wplus <a href="#">v9.0</a> and <a href="#">NRT v9.3</a> and <a href="#">obspack co 2 466 GVe u v9.2 202 40502</a>	Hourly resolution (well-mixed conditions) obspack GLOBALVIE Wplus v8.0 and <a href="#">v9.0 and NRT v9.2</a>	Hourly resolution (well-mixed conditions) obspack GLOBALVIE Wplus <a href="#">v9.0</a> and <a href="#">NRT v9.3</a>	Hourly resolution (well-mixed conditions) obspack GLOBALVIE Wplus <a href="#">v7.0</a> and <a href="#">NRT v9.2</a>	ACOS- GOSAT <a href="#">B9</a> and <a href="#">OCO-2</a> V11.1 and obspack GLOBALVIE Wplus <a href="#">v9.1</a>	OCO-2 ACOS retrievals from NASA, v11.1	OCO-2 v11r data that scaled to WMO 2019 standard	<a href="#">Hourly resolution (well-mixed conditions) obspack GLOBALVIE Wplus v9.1 and NRT v9.2 And OCO-2 b11.1 LN LG</a>	<a href="#">ACOS v11 OCO-2 XCO2 retrievals, scaled to WMO 2019 standard</a>	<a href="#">Hourly resolution (well-mixed conditions) obspack GLOBALVIE Wplus v9.1 and NRT v9.2</a>	Hourly resolution (well-mixed conditions) obspack GLOBALVIE Wplus <a href="#">v9.1</a> and NRT <a href="#">v9.2</a>	Hourly resolution (well-mixed conditions) obspack GLOBALVIE Wplus <a href="#">v9.1</a> and NRT <a href="#">v9.3</a> and <a href="#">JMA</a>	Hourly resolution (well-mixed conditions) obspack GLOBALVIEW plus <a href="#">v9.1</a> and <a href="#">NRT v9.2</a> and <a href="#">GOSAT XCO2 data</a> <a href="#">NIES Level 2</a> <a href="#">product v02.95 and v03.05</a>
Period covered	<a href="#">1976-2023</a>	<a href="#">1979-2023</a>	<a href="#">2001-2023</a>	<a href="#">1990-2023</a>	<a href="#">2000-2023</a>	<a href="#">2010-2023</a>	<a href="#">2015-2023</a>	<a href="#">2015-2023</a>	<a href="#">2015-2023</a>	<a href="#">2015-2023</a>	<a href="#">2001-2023</a>	<a href="#">2001-2023</a>	<a href="#">2001-2023</a>	<a href="#">2010-2023</a>
Prior fluxes														

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Biosphere and fires	Zero	ORCHIDEE-GFEDv4.1s	SIB4-MERRA and GFAS	VISIT and GFEDv4.1s	GFED-CASA and GFED_CMS/ Climatology for the CT-NRT of CT2022 plus statistical flux anomaly model)	CARDAMOM	ORCHIDEE-GFEDv4.1s	ORCHIDEE-MICT and GFEDv4.1s	VEGAS + GFAS	BEPS	CASA v1.0, climatology after 2016 and GFED4.0	CASA v1.0, climatology after 2016 and GFED4.0	CASA-3h	Zeng et al. 2020 and GFAS	Deleted: SIB4.2 and GFEDv4.1s
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Ocean	CarboScope pc v2024E	CMEMS-LSCE-FFNN 2024	CarboScope v2022 and v2023	JMA global ocean mapping (Iida et al., 2021)	Ocean inversion fluxes, Takahashi pCO2	ECCO-Darwin and MOM6	CMEMS-LSCE-FFNN 2023	Takahashi climatology	Jena OC-v2023	JMA Ocean CO2 Map v2023 (Global) and v2024 (regional)	Takahashi climatology	Takahashi climatology	Takahashi climatology	Zeng et al. 2014	Formatted ... [439]
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Fossil fuels (c)	GridFED v2024.0	GridFED 2023.1 with an extrapolation to 2023-24 based on Carbonmonitor and NO2	GridFED 2023.1 and 2024.0	GridFED v2024.0	Miller/CT and ODIAC/NAS A	GridFED v2024.0	GridFED 2023.1 with an extrapolation to 2023-24 based on Carbonmonitor and NO2	GridFED 2024.0	GridFEDv2023.1 and v2024.0	GridFEDv2024.0	GridFED 2024.0	GridFED 2024.0	GridFEDv2024.0	GridFEDv2024.0	Formatted ... [437]
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Transport and optimization															Formatted ... [446]
Transport model	TM3	LMZ v6	TM5	NICAM-TM	TM5	GEOS-CHEM	LMZ v6	GEOS-Chem v12.9.3	GEOS-CHEM v13.0.2	MOZART-4	GEOS-CHEM v12.5	GEOS-CHEM v12.5	MIROC-ACTM	NIES-FLEXART	Formatted ... [452]
Weather forcing	ERA	ECMWF	ECMWF	JRA55	ERA5	MERRA2	ECMWF	MERRA2	MERRA-2	GEOSS	MERRA	MERRA	JRA-55	ERA5/ECMWF/MIROC/ACTM/FLEXART	Formatted ... [459]
Horizontal Resolution	Global 3.83°x5°	global ~90 km in the horizontal	Global 3°x2° Europe 1°x1°, North	global-5 (~223 km)	Global 3°x2° North America	Global 4°x5°	global ~90 km in the horizontal	Global 2°x2.5°	2°x2.5°	2.5°x1.875°	Global 2°x2.5°	Global 4°x5°	2.8°x2.8°	NIES-TM/3.75x3.75 and d FLEXART	Formatted ... [455]
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Optimization	Conjugate gradient (re- ortho- normalization)	Variational	5-week ensemble Kalman smoother	Variational	12-week ensemble Kalman smoother	Variational	Variational	Nonlinear least squares four- dimensional variation (NLS-4DVar)	Ensemble Kalman Filter (LETKF with CEnKF/AAP O)	Ensemble Kalman filter	Ensemble Kalman filter	Ensemble Kalman filter	Bayesian inversion similar to that of Rayner et al. (Tellus 1999)	Variational M10N3

(a) Schuldt et al. 2023. Multi-laboratory compilation of atmospheric carbon dioxide data for the period 1957-2022; obspack\_co2\_1 GLOBALVIEWplus\_v9.0\_2023-09-09; NOAA Earth System Research Laboratory, Global Monitoring Laboratory. <http://doi.org/10.25925/20230801>

(b) Schuldt et al. 2024. Multi-laboratory compilation of atmospheric carbon dioxide data for the period 2023-2024; obspack\_co2\_1\_NRT\_v9.2\_2024-03-25; NOAA Earth System Research Laboratory, Global Monitoring Laboratory. <http://doi.org/10.25925/20240215>

(c) GCP-GridFED v2024.0 and v2023.1 (Jones et al., 2024, 2023) are updates through the year 2023 of the GCP-GridFED dataset presented by Jones et al. (2021b)

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**Table S5: Overview of the Earth System Models (ESMs) and the simulations.**

Model	CanESM5	EC-Earth3-CC	JPSL-CM6A-CO2-LR	MIROC-ES2L	MPI-ESM1-2-LR
Resolution Atmosphere	T63, 49 hybrid levels up to 1hPa	T255, 91 levels	2.5°x1.25°, 79 levels	T42, 40 levels	T63, 47 levels
Resolution Ocean	1° refined meridionally to 1/3° near Equator, 45 levels	1°, 75 levels	1° (nominal), 75 levels	Tripolar (~1°), 62 levels	1.5°, 40 levels
Assimilation Atmosphere	ERA-Interim (Dee et al. 2011) from 1980 to 2018 and ERA5 (Hersbach et al. 2020) afterwards: full-field nudging of temperature, horizontal wind and specific humidity	ERA5 (Hersbach et al. 2020) full-field	None	3D full field wind and T of JRA55 (Kobayashi et al. 2015) with the simplified IAU (Tatebe et al. 2012)	ERA-40 (Uppala et al. 2005) before 1979 and ERA5 (Hersbach et al. 2020) from 1980: Vorticity, divergence, log(p), T; full field with nudging

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<a href="#">Assimilation Ocean</a>	<a href="#">Nudging to 3D potential temperature and salinity from ORAS5 reanalysis (Zuo et al. 2019). Sea surface temperature relaxed to interpolated values from NOAA's OISSTv2 from Nov. 1981 to present, and NOAA's ERSSTv3 prior (Smith et al. 2008).</a>	<a href="#">EN4 (Good et al. 2013) 3D nudging T and S with weaker nudging band around equator. SST and SSS restoring to ORAS5 (Zuo et al. 2019). Atmospheric forcing: DFS5.2 1958-1979 and ERA5 after 1980.</a>	<a href="#">Nudging towards SST (ERSSTv5) and SSS (EN4) using a restoring coefficient dependent on the mixed layer depth (Ortega et al. 2017)</a>	<a href="#">3D full field T, S, and sea-ice concentration of an ocean objective analysis (Ishii and Kimoto 2009) with the simplified IAU (Tatebe et al. 2012)</a>	<a href="#">EN4 (Good et al. 2013) 3D full field T and S with ensemble Kalman filter (Brune et al. 2018)</a>
<a href="#">Assimilation Land</a>	<a href="#">Indirectly through response of CLASS-CTEM to the data-constrained coupled ESM</a>	<a href="#">LPJ-GUESS forced offline with ERA5 1979-2020 after preindustrial spinup+transient up to 1979</a>	<a href="#">None</a>	<a href="#">None</a>	<a href="#">Indirectly initialized by atmospheric and oceanic data assimilation within the fully coupled ESM</a>
<a href="#">Ensemble Size</a>	<a href="#">10</a>	<a href="#">10</a>	<a href="#">10</a>	<a href="#">10</a>	<a href="#">10</a>
<a href="#">Period of reconstruction</a>	<a href="#">1960-2023</a>	<a href="#">1980-2023</a>	<a href="#">1960-2023</a>	<a href="#">1960-2023</a>	<a href="#">1960-2023</a>
<a href="#">Hindcasts and forecasts</a>	<a href="#">1 year starting from Jan. 1st 1980-2024</a>	<a href="#">14 months starting from Nov.1st 1980-2023</a>	<a href="#">1 year starting from Jan. 1st 1981-2024</a>	<a href="#">14 months starting from Nov. 1st 1980-2023</a>	<a href="#">14 months starting from Nov.1st 1980-2023</a>
<a href="#">External forcings</a>	<a href="#">The Coupled Model Intercomparison Project Phase 6 (CMIP6) historical (1960-2014) plus SSP2-4.5 baseline and CovidMIP two year blip scenario (after 2015) (Eyring et al. 2016; Lamboll et al. 2021). The CO2 emissions forcing from 2015 onward are substituted by GCP-GridFED (Jones et al. 2021; 2023) for all the models except for IPSL-CM6A-CO2-LR. Note the difference in global integrated CO2 emissions between CMIP6 CovidMIP and GCP-GridFED in recent years is within the emission uncertainty.</a>				

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<a href="#">References</a>	<a href="#">Swart et al. 2019; Sospedra-Alfonso et al. 2021</a>	<a href="#">Döscher et al. 2021; Bilbao et al., 2021; Bernardello et al., 2024</a>	<a href="#">Boucher et al. 2020</a>	<a href="#">Watanabe et al. 2020</a>	<a href="#">Mauritsen et al. 2019; Li et al. 2023</a>
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**Table S6. Comparison of the projection with realised fossil CO2 emissions (EFOS). The 'Actual' values are first the estimate available using actual data, and the 'Projected' values refers to estimates made before the end of the year for each publication. Projections based on a different method from that described here during 2008-2014 are available in Le Quéré et al. (2016). All values are adjusted for leap years.**

	World		China		USA		EU28 / EU27 (i)		India		Rest of World (ii)	
	Projected	Actual	Projected	Actual	Projected	Actual	Projected	Actual	Projected	Actual	Projected	Actual
2015 (a)			-0.6%	0.06%	-3.9%	-0.7%	-1.5%				1.2%	1.2%
			(-1.6 to 0.5)		(-4.6 to -1.1)		(-5.5 to 0.3)	-2.5%	=	=	=	(-0.2 to 2.6)
2016 (b)	-0.2%	0.20%	-0.5%	-0.3%	-1.7%	-2.1%	=	=	=	=	1.0%	1.5%
			(-1.0 to +1.8)		(-3.8 to +1.3)		(-4.0 to +0.6)				(-0.4 to +2.5)	
2017 (c)			2.0%	1.6%	3.5%	1.5%	-0.4%	-0.5%		2.00%	1.6%	1.6%
			(+0.8 to +3.0)		(+0.7 to +5.4)		(-2.7 to +1.0)		=	(+0.2 to +3.8)	3.9%	(0.0 to +3.2)
2018 (d)			2.7%	2.1%	4.7%	2.3%	2.5%	2.8%	-0.7%	-2.1%	6.3%	8.0%
			(+1.8 to +3.7)		(+2.0 to +7.4)		(+0.5 to +4.5)	(-2.6 to +1.3)		(+4.3 to +8.3)	(+0.5 to +3.0)	
2019 (e)	0.5%	0.1%	2.6%		-2.4%		-1.7%	-4.3%	1.8%	1.0%	0.5%	0.5%
			(-0.3 to +1.4)		(+0.7 to +4.4)	2.2%	(-4.7 to -0.1)	-2.6%	(-5.1 to +1.8%)	(-0.7 to +3.7)	(-0.8 to +1.8)	
2020 (f)	-6.7%	-5.4%	-1.7%	1.4%	-12.2%		-11.3%	-10.6%	-9.1%	-7.3%	-7.4%	-7.4%
							(EU27)					

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2021 (g)	4.8%	5.1%	4.3%	3.5%	6.8%	6.2%	6.3%	6.8%	11.2%	11.1%	3.2%	4.5%
	(4.2% to 5.4%)		(3.0% to 5.4%)		(6.6% to 7.0%)		(4.3% to 8.3%)		(10.7% to 11.7%)		(2.0% to 4.3%)	
2022 (h)	1.1%	0.9%	-1.5%	0.9%	1.6%		-1.0%		5.6%		2.5%	
	(0% to 1.7%)		(-3.0% to 0.1%)		(-0.9% to 4.1%)	1.0%	(-2.9% to 1.0%)	-1.9%	(3.5% to 7.7%)	5.8%	(0.1% to 2.3%)	0.6%
2023 (i)	1.1%	1.3%	4.0%		-3.0%		-7.4%	-8.4%	8.2%		0.4%	
	(0.0% to 2.1%)		(1.9% to 6.1%)	4.9%	(-5% to -1%)	-3.3%	(-9.9% to -4.9%)		(6.7% to 9.7%)	8.2%	(-1.4% to 2.3%)	0.4%
2024 (k)	0.8%		0.2%		-0.6%		-3.8%		4.6%		1.6%	
	(-0.3% to 1.9%)		(-1.6% to 2.0%)		(-2.9% to 1.7%)		(-6.2% to -1.4%)		(3.0% to 6.1%)		(-0.4% to 3.6%)	

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**Table S7 Attribution of fCO<sub>2</sub> measurements for the year 2023 included in SOCATv2024 (Bakker et al., 2016, 2024) to inform ocean fCO<sub>2</sub>-based data products.**

Platform Name	Regions	No. of measurements	Principal Investigators	No. of datasets	Platform Type
Atlantic Explorer	North Atlantic, coastal	48,596	Bates, N. R.; Enright M.	20	Ship
Atlantic Sail	North Atlantic, coastal	16,770	Steinhoff, T.; Körtzinger, A.	3	Ship
Bell M. Shimada	North Pacific, Tropical Pacific, coastal	35,730	Alin, S.; Feely, R.	7	Ship
Cap San Lorenzo	Tropical Atlantic, coastal	18,343	Lefèvre, N.	1	Ship
CCE1 122W 3N	Coastal	1,426	Sutton, A.; Send, U.; Ohman, M.	1	Mooring
CCE2 121W 4N	Coastal	417	Sutton, A.; Send, U.; Ohman, M.	1	Mooring
Colibri	North Atlantic, coastal	24,528	Lefèvre, N.	3	Ship
Equinox	North Atlantic, Tropical Atlantic, coastal	19,612	Wanninkhof, R.; Pierrot, D.	12	Ship
F.G. Walton Smith	Coastal	3,831	Barbero L.; Pierrot, D.; Wanninkhof, R.	3	Ship
Finnmaid	Coastal	311,468	Rehder, G.; Bittig, H. C.; Glockzin, M.	10	Ship
G.O. Sars	Arctic, North Atlantic, coastal	103,965	Skjelvan, I.	12	Ship
GAKOA 149W 60N	Coastal	470	Monacci, N.	1	Mooring
Gordon Gunter	North Atlantic, coastal	24,848	Wanninkhof, R.; Pierrot, D.	4	Ship
Henry B. Bigelow	Coastal	18,661	Wanninkhof, R.; Pierrot, D.	3	Ship
Heron Island	Coastal	1,322	Tilbrook, B.; van Ooijen E.	1	Mooring
Investigator	Southern Ocean	152,788	Tilbrook, B.; Akl, J.; Neill, C.	7	Ship
Kangaroo Island	Southern Ocean	378	Tilbrook, B.; van Ooijen E.	1	Mooring
KC BUOY	Coastal	3,020	Evans, W.	1	Mooring
Keifu Maru II	North Pacific, Tropical Pacific, coastal	7,300	Enyo, K.	5	Ship
Maria Island	Southern Ocean	1,640	Tilbrook, B.; van Ooijen E.	1	Mooring
Marion					
Dufresne	Indian Ocean, Southern Ocean	5,662	Lo Monaco, C.; Metzl, N.	1	Ship
New Century 2	North Atlantic, North Pacific, Tropical Pacific, Southern Ocean, coastal	258,209	Nakaoka, S.-I.; Takao, S.; Sutton, A.; Cronin, M.; Emerson, S.	16	Ship
Papa 145W 50N	North Pacific	820	Emerson, S.	1	Mooring
Quadra Island Field Station	Coastal	78,466	Evans, W.	1	Mooring
R/V Belgica	Coastal	4,485	Theetaert, H.; Gkritzalis, T.	1	Ship
Roger Revelle	Tropical Pacific, Southern Ocean	37,941	Alin, S.; Woosley R. J.; Feely, R.; Martz T. R.	3	Ship

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Ryofu Maru III	North Pacific, Tropical Pacific, coastal	7,454	Enyo, K.	7	Ship
SA Agulhas II	Southern Ocean	7,123	Hamnca, S.; Tsanwani, M.; Monteiro, P. M. S.	1	Ship
Sea Explorer	Southern Ocean, Coastal, Tropical Atlantic, North Atlantic	69,377	Olivier, L.; Landschützer, P.	3	Ship
Seaspan Royal	Coastal	230,720	Evans, W.	6	Ship
Simon Stevin	Coastal	80,488	Gkritzalis, T.; Theetaert, H.; T'Jampens, M.	11	Ship
Soyo Maru	North Pacific, coastal	42,169	Ono, T.	2	Ship
Statsraad Lehmkuhl	North Atlantic, Tropical Atlantic, Southern Ocean, coastal	27,582	Becker, M.; Olsen, A.	2	Ship
Tangaroa	Southern Ocean	15,315	Currie, K. I.	3	Ship
TAO170W ON	Tropical Pacific	2,091	Sutton, A.	1	Mooring
Thomas G. Thompson	North Pacific, Tropical Pacific, Southern Ocean, coastal	29,782	Alin, S.; Feely, R.	5	Ship
Trans Future 5	North Pacific, Tropical Pacific, Southern Ocean, coastal	159,856	Nakaoka, S.-I.; Takao, S.	14	Ship
Tukuma Arctica	North Atlantic, coastal	53,130	Becker, M.; Olsen, A.	17	Ship
Victor					
Angelescu	Southern Ocean	23,904	Berghoff C.; Arbilla L.; Vecchia M.	3	Ship
Wakataka					
Maru	North Pacific, coastal	62,156	Tadokoro, K.; Ono, T.	5	Ship
WHOTS 158W			Sutton, A.; Weller, B.		
23N	Tropical Pacific	1,440	Plueddemann, A.	1	Mooring

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	Alpha Jet Atmospheric eXperiment (AJAX)							
AJAX								
ALF	Alta Floresta		Gatti, L.V.; Gloor, E.; Miller, J.B.;					
A O A			Aircraft Observation of Atmospheric trace gases by JMA				ghq_obs@ met.kishou .go.jp	
BGI	Bradgate, Iowa		Sweeney, C.; Dlugokencky, E.J.					
BNE	Beaver Crossing, Nebraska		Sweeney, C.; Dlugokencky, E.J.					
BRZ	Berezorechka, Russia		Sasakama, N.; Machida, T.					
CAR	Briggsdale, Colorado		Sweeney, C.; Dlugokencky, E.J.					
CMA	Cape May, New Jersey		Sweeney, C.; Dlugokencky, E.J.					
CON	CONTRAIL (Comprehensive Observation Network for TRace gases by AirLiner)	<a href="http://dx.doi.org/10.17595/20180208.001">http://dx.doi.org/10.17595/20180208.001</a>	Machida, T.; Ishijima,					

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<a href="#">CRV</a>	<a href="#">Carbon in Arctic Reservoirs Vulnerability Experiment (CARVE)</a>		<a href="#">Sweeney, C.;</a> <a href="#">Karion, A.;</a> <a href="#">Miller, J.B.;</a> <a href="#">Miller, C.E.;</a> <a href="#">Dlugokencky, E.J.</a>						
<a href="#">DND</a>	<a href="#">Dahlen, North Dakota</a>		<a href="#">Sweeney, C.;</a> <a href="#">Dlugokencky, E.J.</a>						
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<a href="#">HAA</a>	<a href="#">Molokai Island, Hawaii</a>		<a href="#">Sweeney, C.;</a> <a href="#">Dlugokencky, E.J.</a>						
<a href="#">HFM</a>	<a href="#">Harvard University Aircraft Campaign</a>		<a href="#">Wofsy, S.C.</a>						
<a href="#">HIL</a>	<a href="#">Homer, Illinois</a>		<a href="#">Sweeney, C.;</a> <a href="#">Dlugokencky, E.J.</a>						
<a href="#">HIP</a>	<a href="#">HIPPO (HIAPER Pole-to-Pole Observations)</a>	<a href="https://doi.org/10.3334/CDIAC/HIPPO_010">https://doi.org/10.3334/CDIAC/HIPPO_010</a>	<a href="#">Wofsy, S.C.;</a> <a href="#">Stephens, B.B.;</a> <a href="#">Elkins, J.W.;</a> <a href="#">Hintsa, E.J.;</a> <a href="#">Moore, F.</a>						
<a href="#">IAGOS-CORE</a>	<a href="#">In-service Aircraft for a Global Observing System</a>		<a href="#">Christoph Gerbig (Max-Planck-Institut für Biogeochemie, Jena)</a>						
<a href="#">INX</a>	<a href="#">INFLUX (Indianapolis Flux Experiment)</a>		<a href="#">Sweeney, C.;</a> <a href="#">Dlugokencky, E.J.;</a> <a href="#">Shepson, P.B.;</a> <a href="#">Turnbull, J.</a>						
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<a href="#">MAN</a>	<a href="#">Manaus, Brazil</a>		<a href="#">Miller, J.B.;</a> <a href="#">Martins, G.A.;</a> <a href="#">de Souza, R.A.F.</a>						
<a href="#">NHA</a>	<a href="#">Offshore Portsmouth, New Hampshire (Isles of Shoals)</a>		<a href="#">Sweeney, C.;</a> <a href="#">Dlugokencky, E.J.</a>						
<a href="#">OIL</a>	<a href="#">Oglesby, Illinois</a>		<a href="#">Sweeney, C.;</a> <a href="#">Dlugokencky, E.J.</a>						
<a href="#">ORC</a>	<a href="#">ORCAS (O2/N2 Ratio and CO2 Airborne Southern Ocean Study)</a>	<a href="https://doi.org/10.5065/D6SB445X">https://doi.org/10.5065/D6SB445X</a>	<a href="#">Stephens, B.B.;</a> <a href="#">Sweeney, C.;</a> <a href="#">McKain, K.;</a> <a href="#">Kort, E.</a>						
<a href="#">PFA</a>	<a href="#">Poker Flat, Alaska</a>		<a href="#">Sweeney, C.;</a> <a href="#">Dlugokencky, E.J.</a>						
<a href="#">RBA-B</a>	<a href="#">Rio Branco</a>		<a href="#">Gatti, L.V.;</a> <a href="#">Gloor, E.;</a> <a href="#">Miller, J.B.</a>						
<a href="#">RTA</a>	<a href="#">Rarotonga</a>		<a href="#">Sweeney, C.;</a> <a href="#">Dlugokencky, E.J.</a>						

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<a href="#">SCA</a>	<a href="#">Charleston, South Carolina</a>		<a href="#">Sweeney, C.; Dlugokencky, E.J.</a>
<a href="#">SGP</a>	<a href="#">Southern Great Plains, Oklahoma</a>		<a href="#">Sweeney, C.; Dlugokencky, E.J.; Biraud, S.</a>
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<a href="#">TGC</a>	<a href="#">Offshore Corpus Christi, Texas</a>		<a href="#">Sweeney, C.; Dlugokencky, E.J.</a>
<a href="#">THD</a>	<a href="#">Trinidad Head, California</a>		<a href="#">Sweeney, C.; Dlugokencky, E.J.</a>
<a href="#">UGD</a>	<a href="#">Kajjansi Airfield, Kampala, Uganda</a>		<a href="#">McKain, K.; Sweeney, C.</a>
<a href="#">ULB</a>	<a href="#">Ulaanbaatar, Mongolia</a>		<a href="#">Sweeney, C.; Dlugokencky, E.J.</a>
<a href="#">WBI</a>	<a href="#">West Branch, Iowa</a>		<a href="#">Sweeney, C.; Dlugokencky, E.J.</a>
			<a href="#">(a) Schuldt et al. 2023. Multi-laboratory compilation of atmospheric carbon dioxide data for the period 1957-2022; obspack_co2_1 GLOBALVIEWplus_v9.0 2023-09-09; NOAA Earth System Research Laboratory, Global Monitoring Laboratory. <a href="http://doi.org/10.25925/20230801">http://doi.org/10.25925/20230801</a></a>
			<a href="#">(b) Schuldt et al. 2024. Multi-laboratory compilation of atmospheric carbon dioxide data for the period 2023-2024; obspack_co2_1 NRT_v9.2 2024-03-25; NOAA Earth System Research Laboratory, Global Monitoring Laboratory. <a href="http://doi.org/10.25925/20240215">http://doi.org/10.25925/20240215</a></a>

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**Table S9.** Main methodological changes in the global carbon budget since first publication. Methodological changes introduced in one year are kept for the following years unless noted. Empty cells mean there were no methodological changes introduced that year.

Publication year	Fossil fuel emissions			LUC emissions	Reservoirs			Uncertainty & other changes
	Global	Country (territorial)	Country (consumption)		Atmosphere	Ocean	Land	
2006 (a)		Split in regions						
2007 (b)				ELUC based on FAO-FRA 2005; constant ELUC for 2006	1959-1979 data from Mauna Loa; data after 1980 from global average	Based on one ocean model tuned to reproduced observed 1990s sink		$\pm 1\sigma$ provided for all components
2008 (c)				Constant ELUC for 2007				
2009 (d)		Split between Annex B and non-Annex B	Results from an independent study discussed	Fire-based emission anomalies used for 2006-2008		Based on four ocean models normalised to observations with constant delta	First use of five DGVMs to compare with budget residual	
2010 (e)	Projection for current year based on GDP	Emissions for top emitters		ELUC updated with FAO-FRA 2010				

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2011 (f)			Split between Annex B and non-Annex B					
2012 (g)		129 countries from 1959	129 countries and regions from 1990-2010 based on GTAP8.0	ELUC for 1997-2011 includes interannual anomalies from fire-based emissions	All years from global average	Based on 5 ocean models normalised to observations with ratio	Ten DGVMs available for SLAND; First use of four models to compare with ELUC	
2013 (h)		250 countriesb	134 countries and regions 1990-2011 based on GTAP8.1, with detailed estimates for years 1997, 2001, 2004, and 2007	ELUC for 2012 estimated from 2001-2010 average		Based on six models compared with two data-products to year 2011	Coordinated DGVM experiments for SLAND and ELUC	Confidence levels; cumulative emissions budget from 1750
2014 (i)	Three years of BP data	Three years of BP data	Extended to 2012 with updated GDP data	ELUC for 1997-2013 includes interannual anomalies from fire-based emissions		Based on seven models	Based on ten models	Inclusion of breakdown of the sinks in three latitude bands and comparison with three atmospheric inversions

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2015 (j)	Projection for current year based Jan-Aug data	National emissions from UNFCCC extended to 2014 also provided	Detailed estimates introduced for 2011 based on GTAP9			Based on eight models	Based on ten models with assessment of minimum realism	The decadal uncertainty for the DGVM ensemble mean now uses $\pm 1\sigma$ of the decadal spread across models	Inserted Cells ... [1272]
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2016 (k)	Two years of BP data	Added three small countries; China's emissions from 1990 from BP data (this release only)		Preliminary ELUC using FRA-2015 shown for comparison; use of five DGVMs		Based on seven models	Based on fourteen models	Discussion of projection for full budget for current year	
2017 (l)	Projection includes India-specific data			Average of two bookkeeping models; use of 12 DGVMs		Based on eight models that match the observed sink for the 1990s; no longer normalised	Based on 15 models that meet observation-based criteria (see Sect. 2.5)	Land multi-model average now used in main carbon budget with the carbon imbalance presented separately; new table of key uncertainties	
2018 (m)	Revision in cement emissions; Projection includes EU-specific data	Aggregation of overseas territories into governing nations for total of 213		Average of two bookkeeping models; use of 16 DGVMs	Use of four atmospheric inversions	Based on seven models	Based on 16 models; revised atmospheric forcing from CRUNCEP to CRUJRA	Introduction of metrics for evaluation of individual models using observations	







**Table S10: Translation of global carbon cycle models' land flux definitions to the definition of the LULUCF net flux used by National GHG Inventory (NGHGI) reports to UNFCCC. Non-intact lands are used here as proxy for "managed lands" in the country reporting. NGHGs are gap-filled (see Sec. C.2.3 for details). For comparison, we provide the net land flux on managed land from atmospheric inversions and FAOSTAT estimates. Units are GtC yr<sup>-1</sup>.**

Carbon flux		Source	2004-2013					2014-2023				
ELUC	Bookkeeping estimates from Table S5		▼					▼	▼	▼	▼	▼
		1.41						1.13				
SLAND total	DGVMs from Table S5		▼	▼	▼	▼	▼	▼	▼	▼	▼	▼
								3.15				3.19
SLAND in non-intact forest	DGVMs							1.75				1.83
ELUC minus SLAND in non-intact forest	Bookkeeping ELUC & DGVM SLAND		▼					▼	▼	▼	▼	▼
		-0.34						0.70				
LULUCF	NGHGs		▼	▼	▼	▼	▼	▼	▼	▼	▼	▼
												-0.57
Net land flux on managed land	Atmospheric inversions		▼	▼	▼	▼	▼	▼	▼	▼	▼	▼
								-0.80				0.69
LULUCF	FAOSTAT											
								0.32				0.30

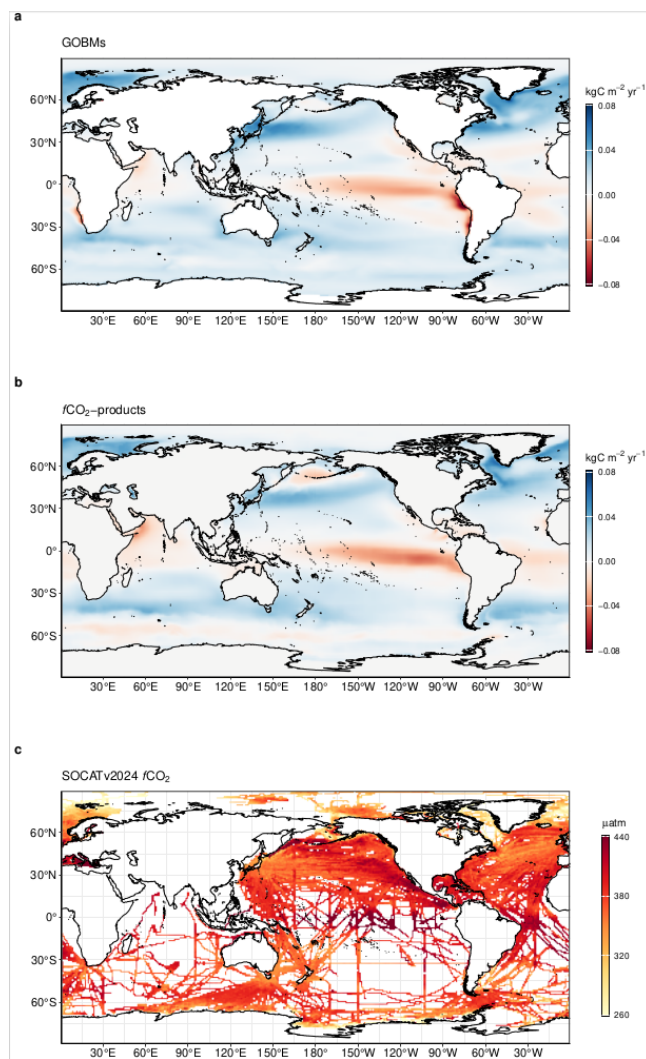
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**Table S11 - Evaluation of global ocean biogeochemistry models based on comparison with observation-based interior ocean carbon accumulation and process-based evaluation metrics for Atlantic Meridional Overturning Circulation (AMOC), Southern Ocean sea surface salinity and surface ocean Revelle factor (following the RECCAP2 ocean model evaluation chapter, Terhaar et al., 2024) and Southern Ocean stratification index (Bourgeois et al., 2022). See supplementary text C3.3 for details of calculation and observational data sources. Note that AMOC from MOM6-Cobalt (Princeton) is only available between 2018 - 2022, which is the value reported here**

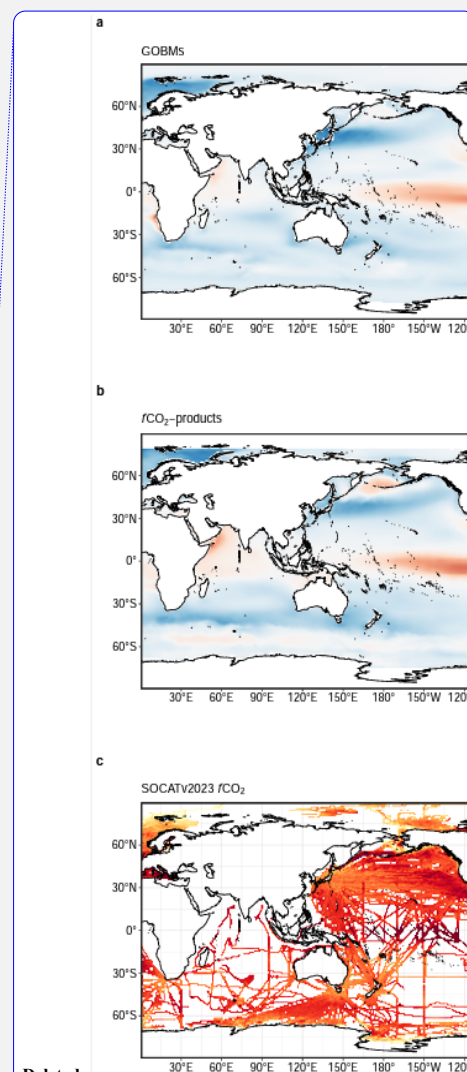
		Global Ocean Biogeochemistry Models									
Metric	Observations	ACCES S (CSIRO)	CESM- ETHZ	FESO M2.1- REco M	MOM 6- Cobalt (Princeton)	MPIO M- HAMO CC6	MRI- ESM2- 3	NEMO - PISCES (IPSL)	NEMO - PlankT OM12	NEMO 3.6- PISCES v2-gas (CNR M)	NorES M- OC1.2
<b>Interior ocean anthropogenic carbon accumulation in GtC yr<sup>-1</sup></b>											
Global (1994-2007, Gruber et al., 2019)	33.8 ± 4.0	36.4	26.0	31.4	27.1	19.9	27.4	28.9	25.4	27.1	33.6
North (1994-2007, Gruber et al., 2019)	5.9	6.4	5.3	5.9	5.1	3.6	5.6	6.0	4.3	5.8	6.8
Tropics (1994-2007, Gruber et al., 2019)	17.5	15.0	8.7	13.3	11.5	9.1	12.5	12.8	12.5	12.5	13.7
South (1994-2007, Gruber et al., 2019)	10.4	15.0	12.0	12.3	10.6	7.2	9.4	10.1	8.6	8.8	12.9
Global (1994-2004, Müller et al., 2023)	29.3 ± 2.5	24.6	19.5	24.1	20.6	15.3	20.3	21.9	18.5	21.2	24.8
Global (2004-2014, Müller et al., 2023)	27.3 ± 2.5	31.4	22.5	27.4	24.2	18.5	23.8	25.0	22.4	23.8	28.5
<b>Atlantic Meridional Overturning Circulation at 26°N, 2005-2022 in Sv (Moat et al., 2023)</b>	17.0 ± 1.3	9.7	13.0	10.2	10.7	15.3	13.5	14.2	17.9	13.1	22.9
<b>Southern Ocean sea surface salinity 2005-2022 in psu (Good et al., 2013)</b>											
subpolar seasonally stratified biome (SPSS)	33.942	34.262	33.809	34.295	34.061	33.925	34.074	34.239	33.873	33.824	34.116
subpolar seasonally stratified and subtropical seasonally stratified biomes (SPSS+STSS)	34.307	34.577	34.185	34.565	34.385	34.254	34.363	34.554	34.358	34.124	34.506
<b>Southern Ocean stratification index 2005-2022, in kg m<sup>-3</sup> (Bourgeois et al., 2022, Good et al., 2013)</b>	5.88	5.45	5.97	5.68	6.13	5.97	6.03	5.60	5.06	6.18	5.76

<i>Surface ocean Revelle factor</i>											
1997-2007, unitless (GLODAPv2.2016, Lauvset et al., 2016)	<a href="#">10.44</a>	<a href="#">10.61</a>	<a href="#">10.33</a>	<a href="#">10.65</a>	<a href="#">10.34</a>	<a href="#">10.72</a>	<a href="#">10.60</a>	<a href="#">10.65</a>	<a href="#">10.49</a>	<a href="#">10.77</a>	<a href="#">10.58</a>
2005-2021, unitless (OceanSODA v2023, updated from Gregor and Gruber, 2021)	<a href="#">10.62</a>	<a href="#">10.77</a>	<a href="#">10.52</a>	<a href="#">10.84</a>	<a href="#">10.52</a>	<a href="#">10.93</a>	<a href="#">10.79</a>	<a href="#">10.81</a>	<a href="#">10.65</a>	<a href="#">10.93</a>	<a href="#">10.75</a>

## Supplementary Figures



**Figure S1.** Ensemble mean air-sea CO<sub>2</sub> flux from a) global ocean biogeochemistry models and b)  $f\text{CO}_2$  based data products, averaged over [2014-2023](#) period ( $\text{kgC m}^{-2} \text{yr}^{-1}$ ). Positive numbers indicate a flux into the ocean. c) gridded SOCAT [v2024](#)  $f\text{CO}_2$  measurements, averaged over the [2014-2023](#) period ( $\mu\text{atm}$ ). In (a) model simulation A is shown. The  $f\text{CO}_2$ -products represent the contemporary flux, i.e. including outgassing of riverine carbon, which is estimated to amount to  $0.65 \text{ GtC yr}^{-1}$  globally.



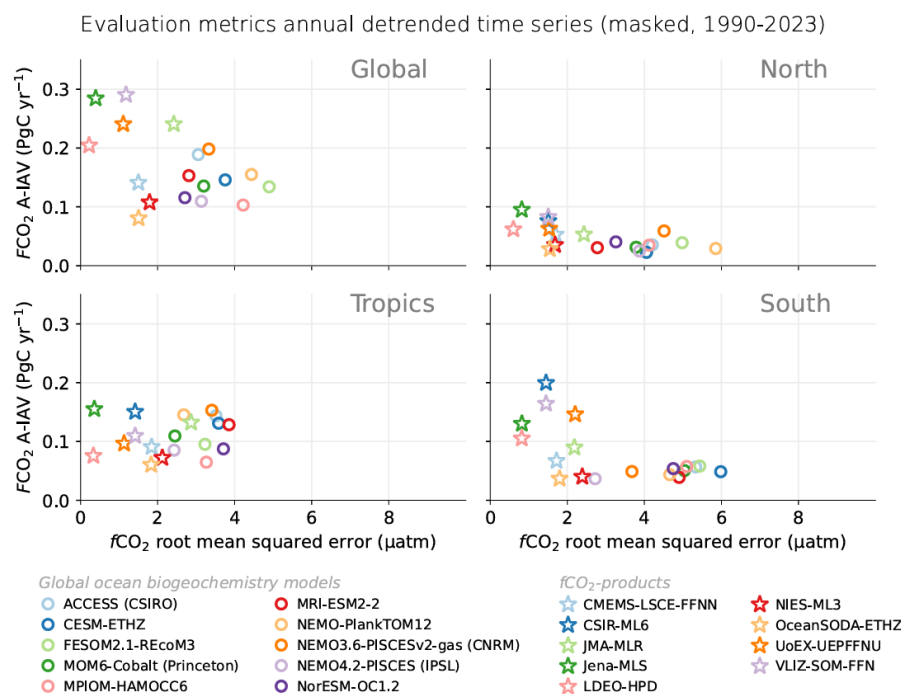
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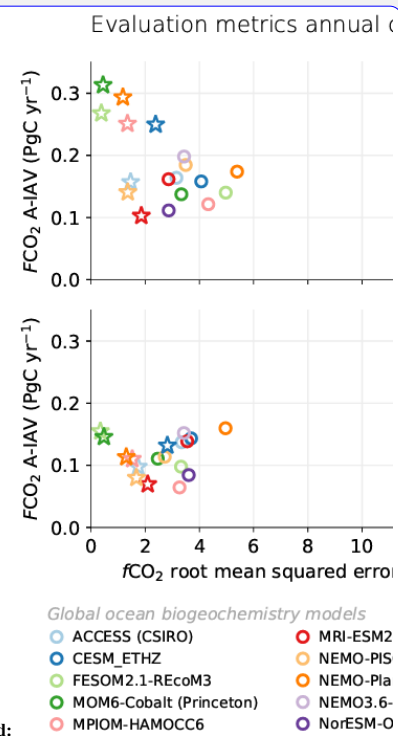
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**Figure S2.** Evaluation of the GOBMs and  $f\text{CO}_2$ -products using the root mean squared error (RMSE) for the period 1990 to 2023, between the individual surface ocean  $f\text{CO}_2$  mapping schemes and the SOCAT v2024 database. The y-axis shows the amplitude of the interannual variability of the air-sea  $\text{CO}_2$  flux (A-IAV, taken as the standard deviation of the detrended annual time series). Results are presented for the globe, north ( $>30^\circ\text{N}$ ), tropics ( $30^\circ\text{S}$ - $30^\circ\text{N}$ ), and south ( $<30^\circ\text{S}$ ) for the GOBMs (see legend, circles) and for the  $f\text{CO}_2$ -based data products (star symbols). The  $f\text{CO}_2$ -products use the SOCAT database and therefore are not independent from the data (see Section 2.5.1).



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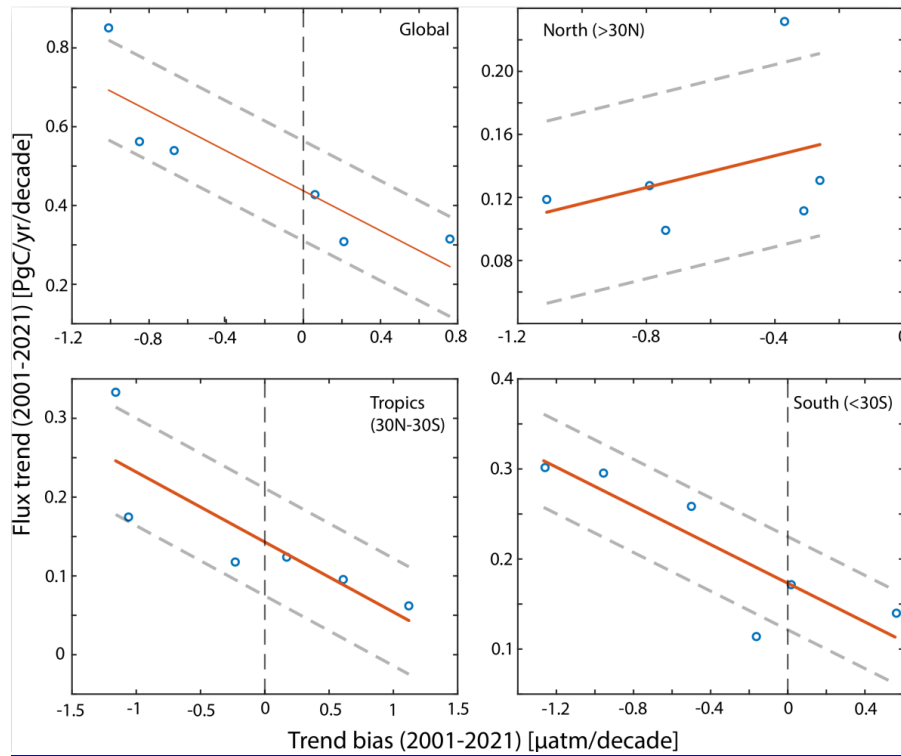
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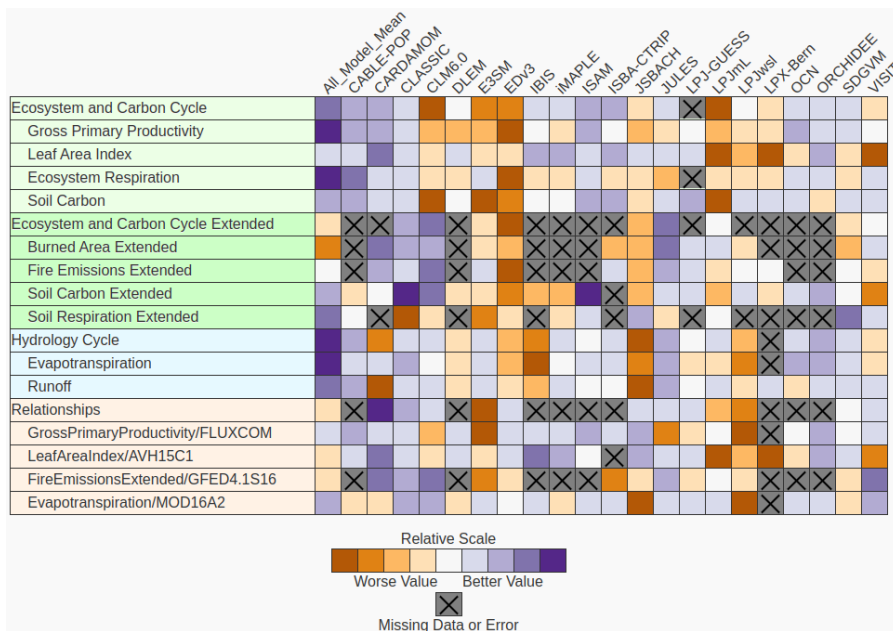
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**Figure S3.** Trend evaluation of six from the eight  $f\text{CO}_2$ -products used for  $\text{SoCEAN}$  (blue circles - CSIR-ML6, NIES-ML3, VLIZ-SOMFFN, OceanSODA-ETHZv2, JMA-MLR, Jena-MLS). The x-axis represents the mean  $f\text{CO}_2$  trend bias from a model subsampling exercise (following Hauck et al., 2023) using four of the GCB2023 GOBMs (CESM, FESOM-REcoM, IPSL and MRI-ESM). The y-axis represents the flux trend as submitted by the  $f\text{CO}_2$  product to this study. Besides the northern hemisphere, where all of the six  $f\text{CO}_2$ -products overestimate the subsampled model trend, there is a clear relationship between the trend reconstruction bias and the flux trend (red line with grey dashed lines representing the 1 sigma uncertainty interval), indicating that flux trends are sensitive to the  $f\text{CO}_2$ -products ability to reconstruction biases.

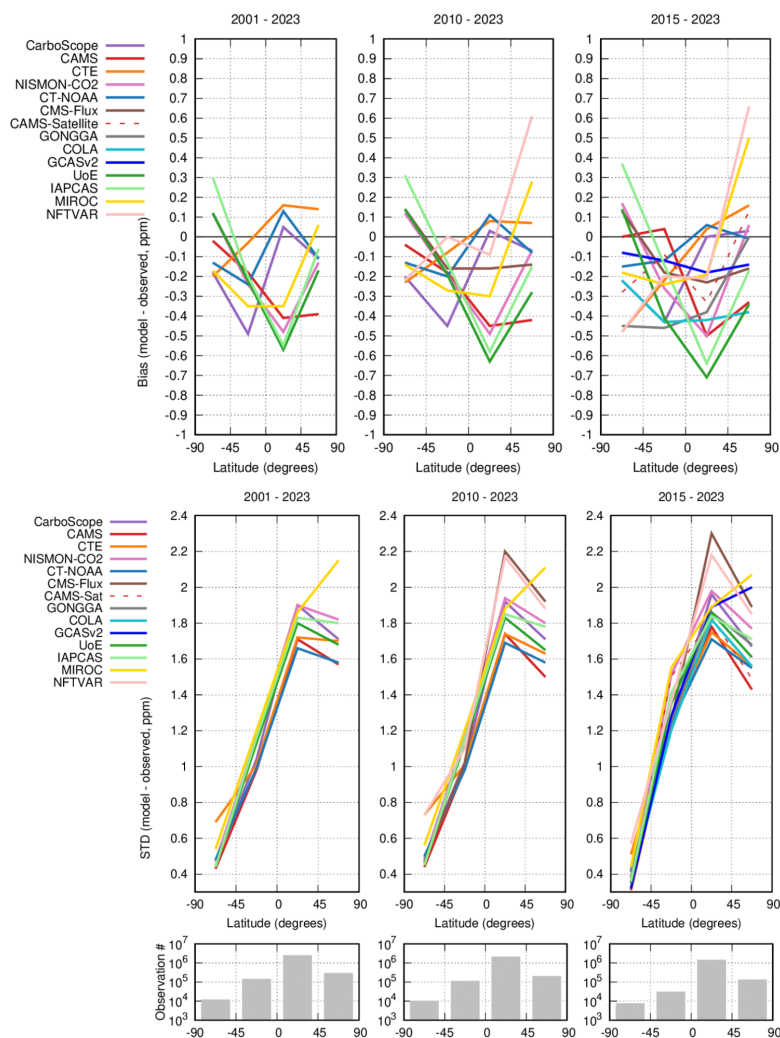


**Figure S4.** Evaluation of the DGVMs using the International Land Model Benchmarking system (ILAMB; Collier et al., 2018) Skill scores relative to other models. The benchmarking is done with observations for GPP and ecosystem respiration (Reichstein et al., 2007; Lasslop et al., 2010; Knauer et al., 2018; Jung et al., 2017; Tramontana et al., 2016; Alemohammad et al., 2017), leaf area index (Vermote, 2019; Claverie et al., 2016; De Kauwe et al., 2011; Myneni et al., 1997), soil carbon (Hugelius et al., 2013; Fischer et al., 2008), evapotranspiration (De Kauwe et al., 2011; Martens et al., 2017; Miralles et al., 2011; Mu et al., 2011), and runoff (Dai and Trenberth, 2002; Hobeichi et al., 2019; Hobeichi et al., 2020). Metrics include relationships between carbon cycle variables, precipitation (Adler et al., 2003) and temperature (Harris et al., 2014). For each model–observation comparison a series of error metrics are calculated, scores are then calculated as an exponential function of each error metric, and finally for each variable the multiple scores from different metrics and observational datasets are combined to give the overall variable scores. Overall variable scores increase from 0 to 1 with improvements in model performance. The set of error metrics vary with dataset and can include metrics based on the period mean, bias, root mean squared error, spatial distribution, interannual variability, and seasonal cycle. The relative skill score shown is a Z score, which indicates in units of standard deviation the model scores relative to the mean score for a given variable. Grey boxes represent missing model data.

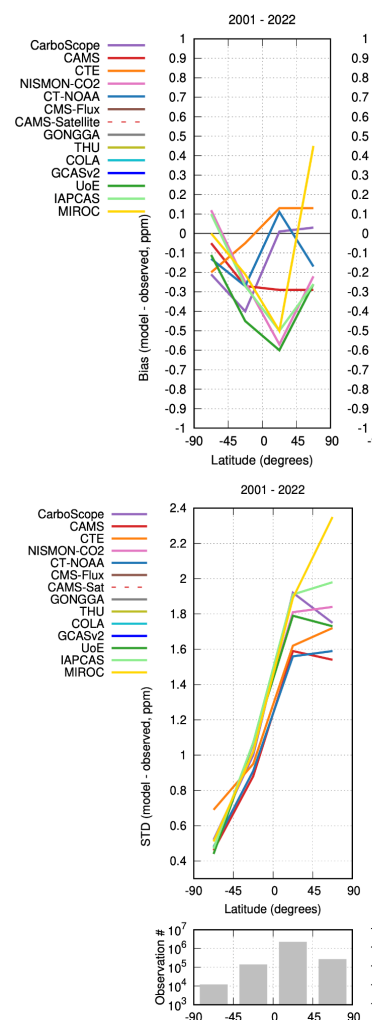
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**Figure S5.** Evaluation of the atmospheric inversion products. The mean of the model minus observations is shown for four latitude bands in three periods: (first panel) 2001–2023, (second panel) 2010–2023, (third panel) 2015–2023. The 14 systems are compared to independent CO<sub>2</sub> observations from aircraft over many places of the world between 2 and 7 km above sea level. Aircraft measurements archived in the Cooperative Global Atmospheric Data Integration Project (Schuldt et al. 2023, Schuldt et al. 2024) from sites, campaigns or programs that have not been assimilated and cover at least 9 months (except for SH programs) between 2001 and 2023, have been used to compute the biases (top row) and their standard deviations (middle row) in four 45° latitude bins. Land and ocean data are used without distinction, and observation density varies strongly with latitude and time as seen on the lower panels.



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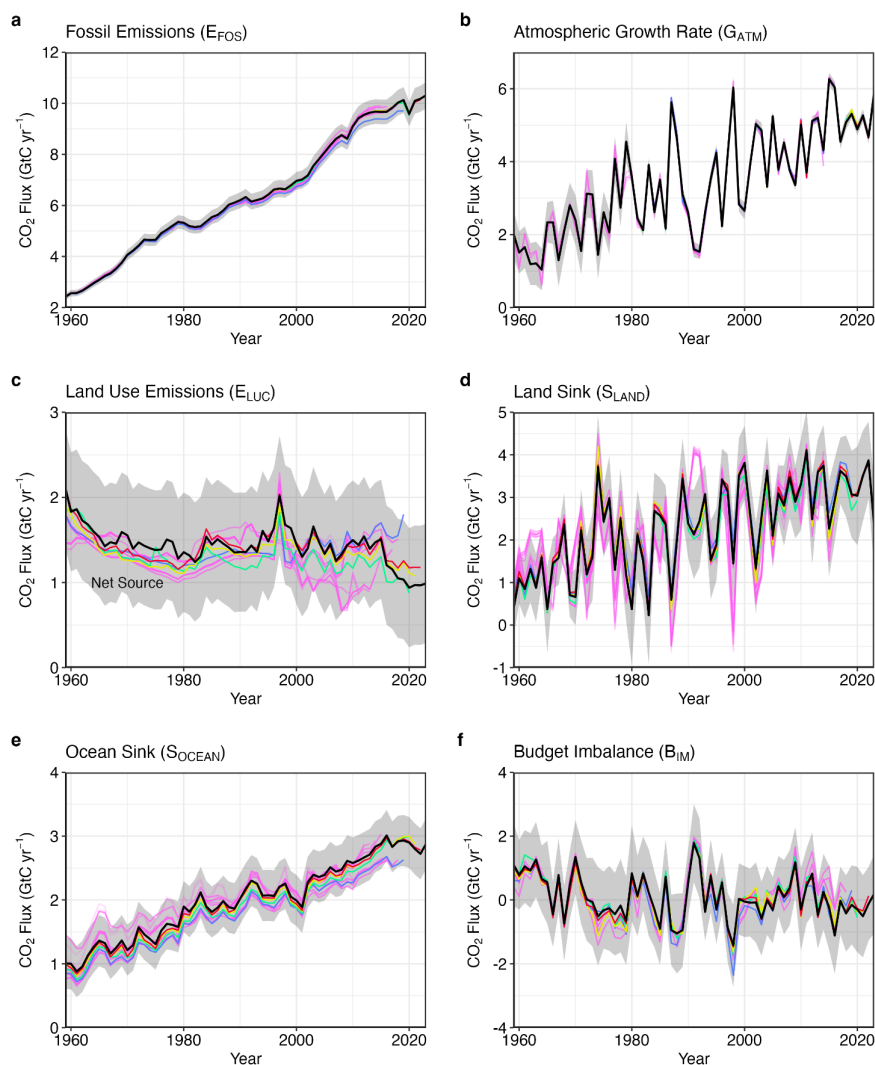
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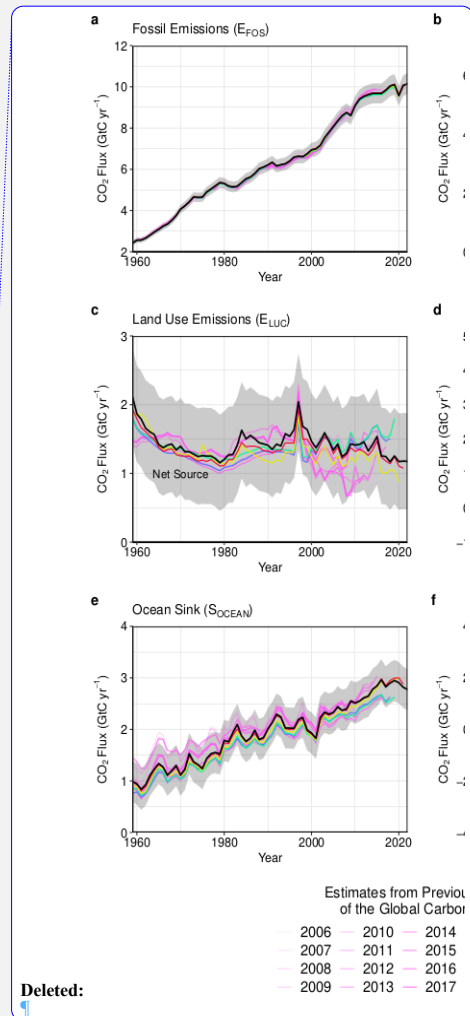
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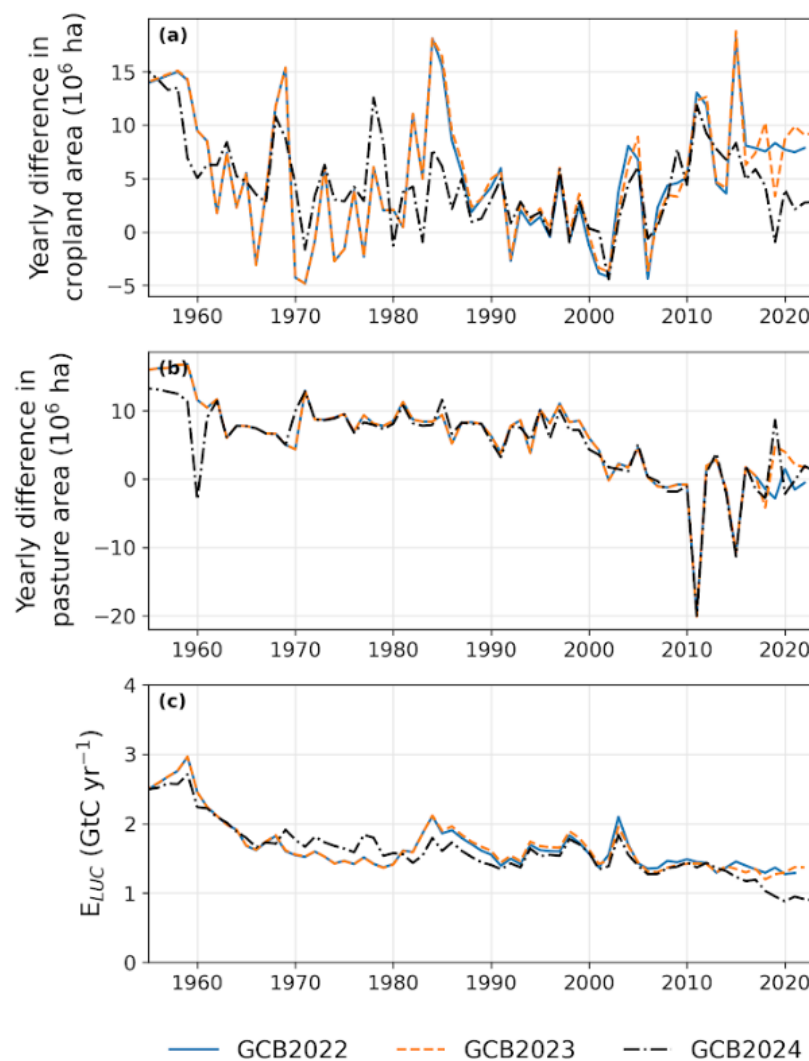
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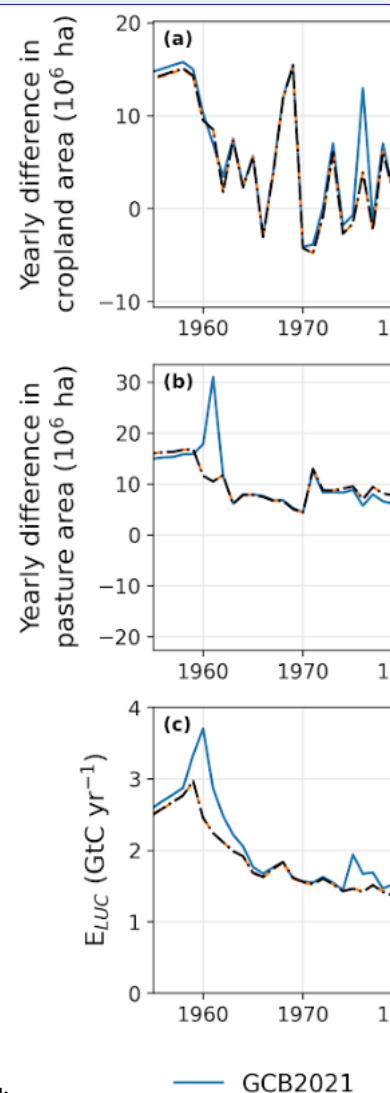
**Figure S6.** Comparison of the estimates of each component of the global carbon budget in this study (black line) with the estimates released annually by the GCP since 2006. Grey shading shows the uncertainty bounds representing  $\pm 1$  standard deviation of the current global carbon budget, based on the uncertainty assessments



described in Supplement S1 to S4. CO<sub>2</sub> emissions from (a) fossil CO<sub>2</sub> emissions excluding cement carbonation ( $E_{FOS}$ ), and (b) land-use change ( $E_{LUC}$ ), as well as their partitioning among (c) the atmosphere ( $G_{ATM}$ ), (d) the land ( $S_{LAND}$ ), and (e) the ocean ( $S_{OCEAN}$ ). See legend for the corresponding years, and Tables 3 and A8 for description of changes in methodology. The budget year corresponds to the year when the budget was first released. All values are in GtC yr<sup>-1</sup>.



**Figure S7.** Differences in the HYDE/LUH2 land-use forcing used for the global carbon budgets GCB2022 (Friedlingstein et al., 2022b), GCB2023 (Friedlingstein et al., 2023), and GCB2024 (this paper). Shown are



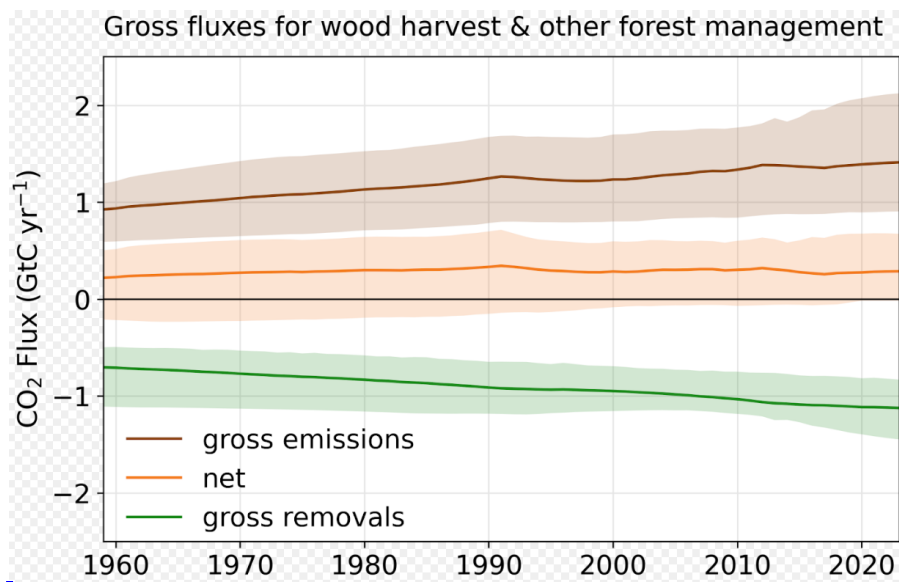
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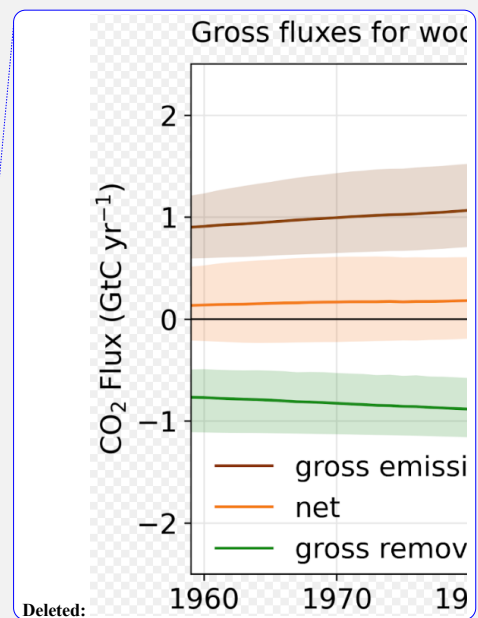
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year-to-year changes in cropland area (top panel) and pasture area (middle panel). To illustrate the relevance of the update in the land-use forcing to the recent trends in  $E_{LUC}$ , the bottom panel shows the land-use emission estimate from the bookkeeping model BLUE (original model output, i.e., excluding emissions from peat fire and peat drainage).



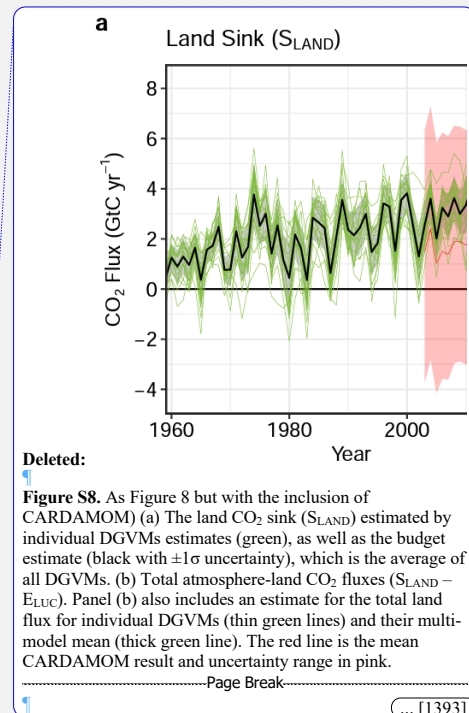
**Figure S8:** Split of net fluxes from wood harvest and other forest management into gross emissions and gross removals. Solid lines denote the average of the three bookkeeping models and shaded areas the full range (min-max) of the bookkeeping model estimates.



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**Figure S9.** Fire carbon emissions for the months January-September for each year 2003-2024 from two global fire emissions products. **(Top row)** Global emissions. **(Middle row)** Emissions for the northern hemisphere extratropics ( $>30^{\circ}$  N), tropics ( $30^{\circ}$  N- $30^{\circ}$  S) and southern extratropics ( $>30^{\circ}$  S). **(Bottom row)** Emissions by RECCAP2 region. The Global Fire Assimilation System (GFAS; Di Giuseppe et al., 2018) (**left column**) and the Global Fire Emissions Database (GFED, version 4.1s; van der Werf et al., 2017) (**right column**) are among the most widely applied global fire emissions products based on satellite remote sensing of fire. GFED relies on the post-fire detection of burned areas combined with fuel consumption factors. GFAS relies on the detection of thermal energy release during active fires.



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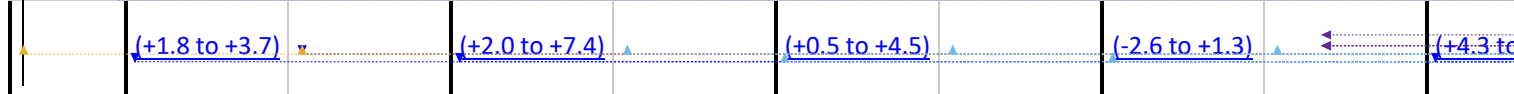
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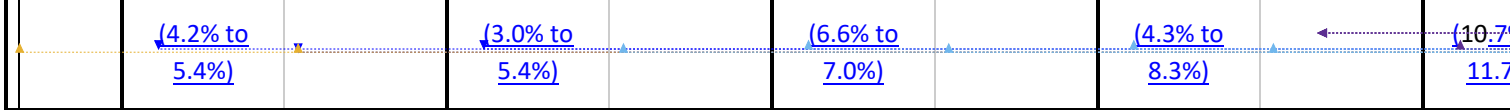
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