

Supplementary Information:

**An in-situ daily dataset for benchmarking temporal variability
of groundwater recharge**

Pragnaditya Malakar^{1,2}, Aatish Anshuman¹, Mukesh Kumar¹, Georgios Boumis¹, T. Prabhakar
5 Clement¹, Arik Tashie¹, Hitesh Thakur¹, Nagaraj Bhat¹, Lokendra Rathore¹

¹Department of Civil, Construction, and Environmental Engineering, University of Alabama, Tuscaloosa, AL,
USA

²Department of Geological Sciences, Jadavpur University, Kolkata, India

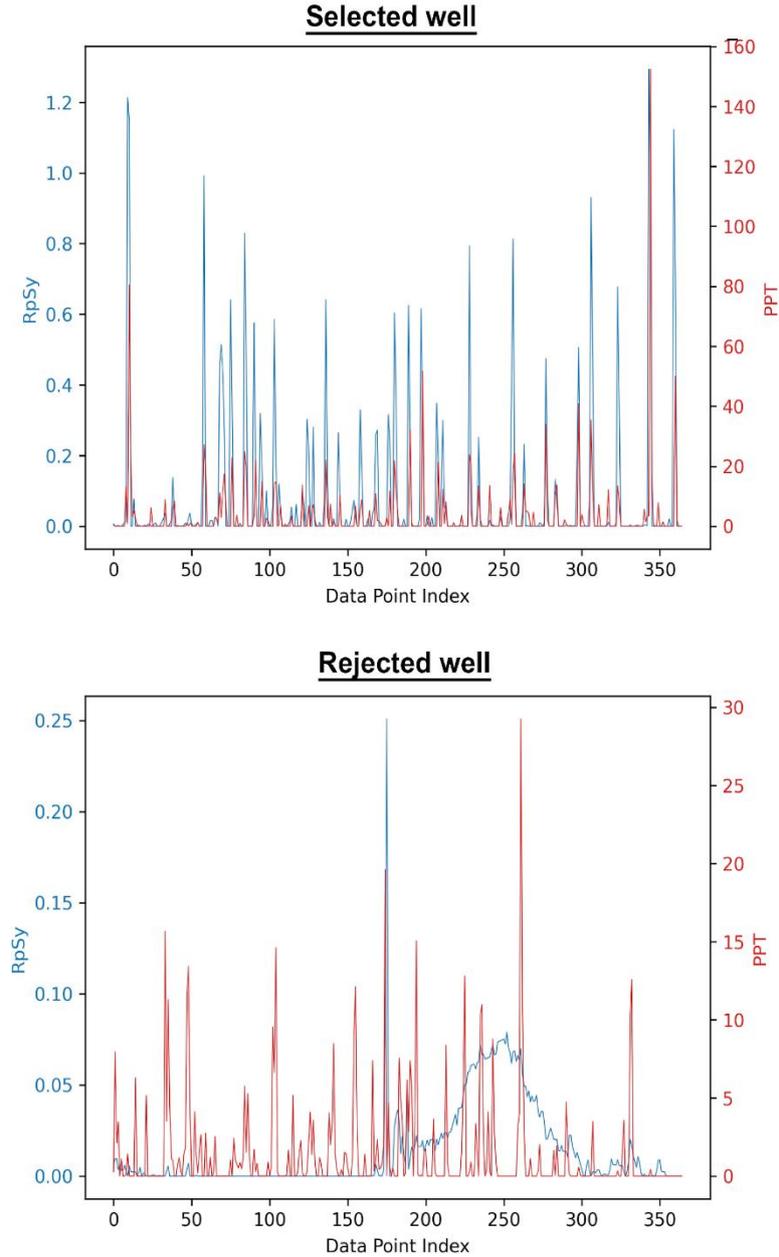
Correspondence to: Mukesh Kumar (mkumar4@eng.ua.edu) and Pragnaditya Malakar
10 (pragnadityamalakar@gmail.com)

Contents of the supplementary information:

15 **Figures S1 to S5**

Table S1

Supplementary text S1:



20 **Figure S1: The hydrographs (a) and (b) show precipitation (Ppt) in mm, and estimated recharge per unit specific yield (RpSy, discussed later) in m per day, for a selected and a rejected well, respectively. The selection/rejection is implemented based on the maximum lag correlation threshold, which ensures that the well is more likely to experience event based GWR response to precipitation signals.**

Supplementary text S1: Comparison of the nearest stream and groundwater level

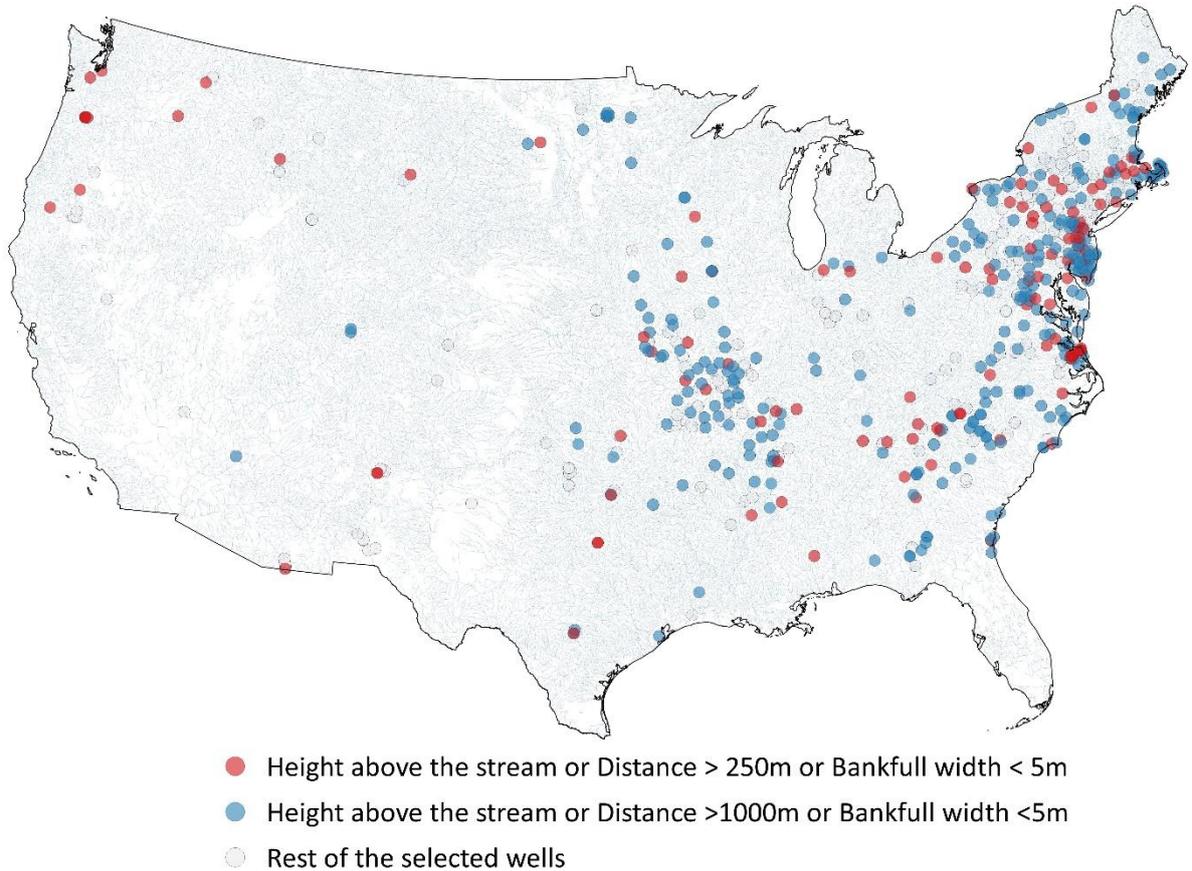
25 Jasechko et al. (2021) compared the elevations of the groundwater level and a constant stream level at bank full height. This study uses a similar approach by identifying the nearest streams to individual wells using high-resolution National Hydrography Dataset Plus Version 2 (McKay et al., 2012). Groundwater well and stream bank elevations are extracted using a 10m resolution 3DEP digital elevation model (DEM) (U.S. Geological Survey, 2019), along with groundwater level time series and bank full-depth estimates (Wieczorek et al., 2019). The elevation difference between groundwater and river levels (E_{diff}) is calculated as follows.

30

$$E_{diff} = E_{GWL} - med(Depth_{GWL}) - E_{Stream} + BFH_{stream} \quad (S1)$$

Here E_{GWL} and E_{Stream} denote groundwater level and river level elevation which are extracted from the DEM. The terms $med(Depth_{GWL})$ and BFH_{stream} represent the median of depth to groundwater level observations and

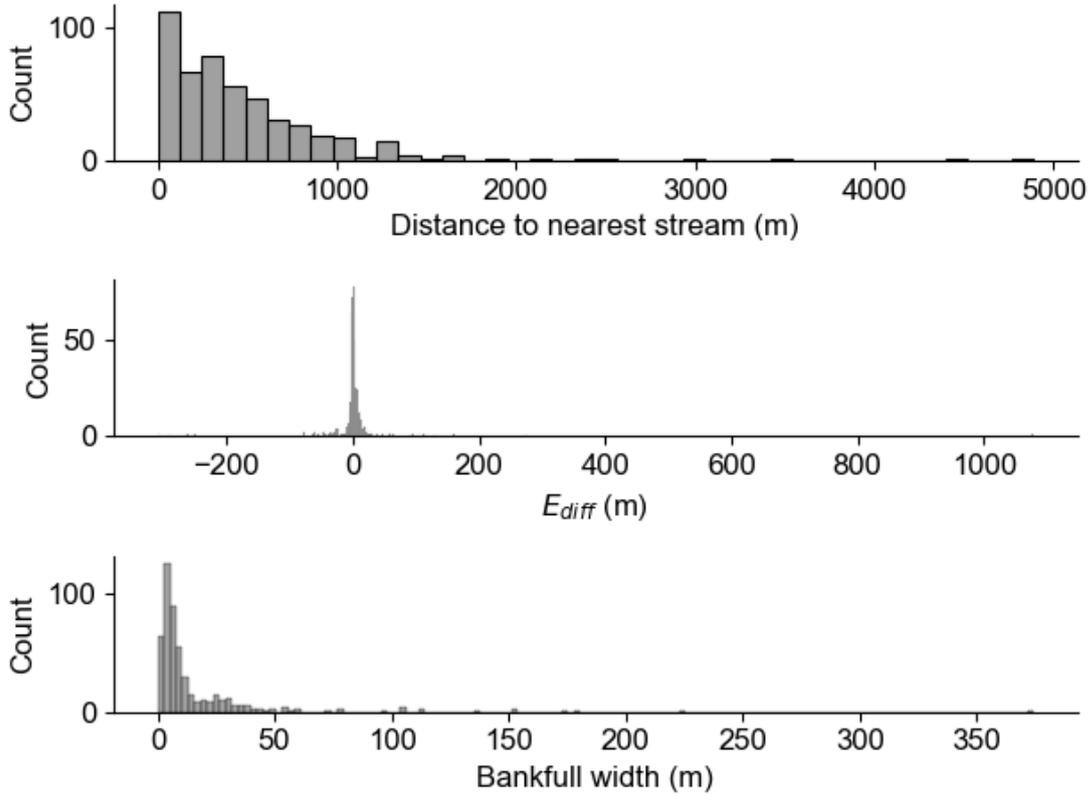
35 bank full height of stream respectively. Eq. (S1) assumes that the elevations extracted using DEM correspond to the elevation of the bank. Since NHD data represents a high-resolution stream network, this assumption is valid due to the likelihood of DEM resolutions being coarser than stream widths. The locations of the wells, distances to nearest streams, differences in elevations, bank full depth and bank full-width estimates are plotted in Figure S2 and S3. Streams with negative E_{diff} values can be classified as losing or influent streams, but this classification also depends on the distance to the nearest well. When the distance is large, the influent flux from stream to well may be negligible. Another important factor is bank full width; this study assumes that rivers with a bank full width of less than 5m are narrow enough to not significantly contribute to groundwater recharge. Considering these criteria, the number of wells used for benchmarking can be narrowed down, as given in Table S1 and Figure S2, S3.



45 **Figure S2: The number of wells selected based on different criteria, which includes distance the nearest stream (shown using blue lines).**

Table S1: The number of wells selected based on different criteria of the nearest stream

Condition of selection of wells	No. of stations
Height above the stream or Distance >250m	388
Height above the stream or Distance >1000m	304
Height above the stream or Distance >250m or Bankfull width <5m	420
Height above the stream or Distance >1000m or Bankfull width <5m	380



50 **Figure S3:** Number of wells with respect to distance to the nearest stream, E_{diff} and bank full width (in meters)

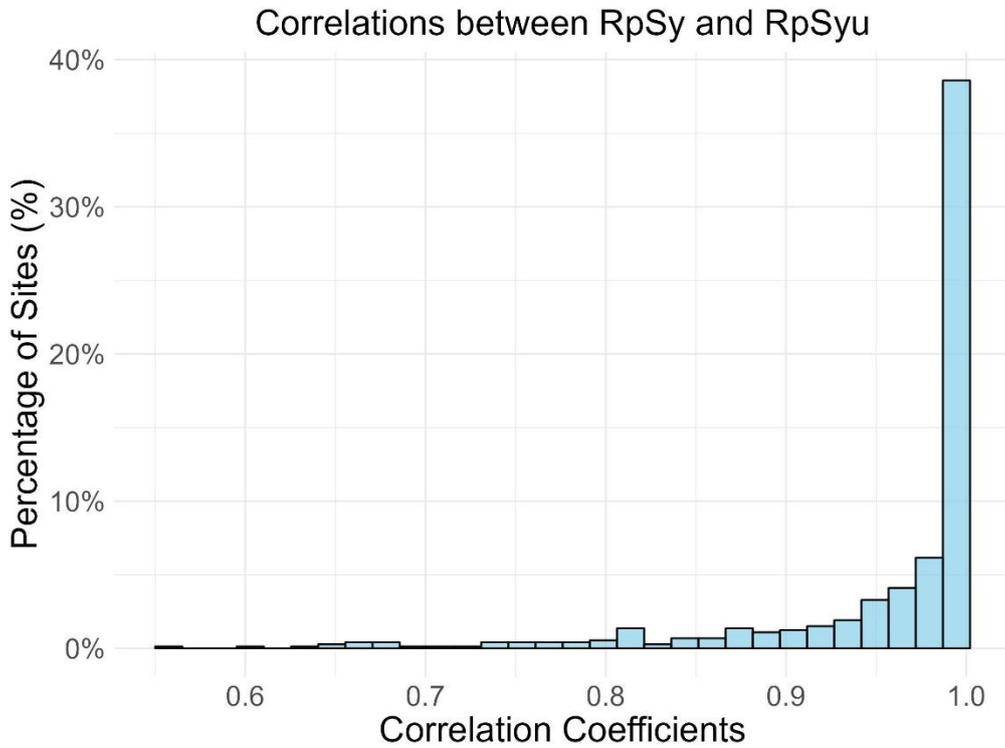


Figure S4: Correlations between RpSy and RpSyu.

RpSy data files for each wells

Date	RpSy (m)	RpSyu (m)
mm ₁ /dd ₁ /yyyy ₁	rpsy ₁₁₁	rpsyu ₁₁₁
mm ₁ /dd ₂ /yyyy ₁	rpsy ₁₂₁	rpsyu ₁₂₁
mm ₁ /dd ₃ /yyyy ₁	rpsy ₁₃₁	rpsyu ₁₃₁
mm ₁ /dd ₄ /yyyy ₁	rpsy ₁₄₁	rpsyu ₁₄₁
mm ₁ /dd ₅ /yyyy ₁	rpsy ₁₅₁	rpsyu ₁₅₁
mm ₁ /dd ₆ /yyyy ₁	rpsy ₁₆₁	rpsyu ₁₆₁
mm ₁ /dd ₇ /yyyy ₁	rpsy ₁₇₁	rpsyu ₁₇₁
.	.	.
.	.	.
.	.	.
mm ₂ /dd ₁ /yyyy ₁	rpsy ₂₁₁	rpsyu ₂₁₁
mm ₂ /dd ₂ /yyyy ₁	rpsy ₂₂₁	rpsyu ₂₂₁
.	.	.
.	.	.
.	.	.
mm ₁ /dd ₁ /yyyy ₂	rpsy ₁₁₂	rpsyu ₁₁₂
.	.	.
.	.	.
.	.	.
mm _p /dd _q /yyyy _r	rpsy _{pqr}	rpsyu _{pqr}

Site information for all selected wells

ID	Lat	Long	Depth (m)
j ₁	x ₁	y ₁	d ₁
j ₂	x ₂	y ₂	d ₂
j ₃	x ₃	y ₃	d ₃
j ₄	x ₄	y ₄	d ₄
j ₅	x ₅	y ₅	d ₅
j ₆	x ₆	y ₆	d ₆
j ₇	x ₇	y ₇	d ₇
.	.	.	.
.	.	.	.
.	.	.	.
j ₈₁	x ₈₁	y ₈₁	d ₈₁

55 **Figure S5: File format of the RpSy and RpSyu dataset.**

References

- Bhanja, S. N., Mukherjee, A., Rangarajan, R., Scanlon, B. R., Malakar, P., & Verma, S. (2019). Long-term groundwater recharge rates across India by in situ measurements. *Hydrology and Earth System Sciences*, 23(2). <https://doi.org/10.5194/hess-23-711-2019>
- 60 Brooks, R., & Corey, A. (1964). Hydraulic properties of porous media. *Hydrology Papers, Colorado State University*, 3 (March).
- Chaney, N. W., Minasny, B., Herman, J. D., Nauman, T. W., Brungard, C. W., Morgan, C. L. S., McBratney, A. B., Wood, E. F., & Yimam, Y. (2019). POLARIS Soil Properties: 30-m Probabilistic Maps of Soil Properties Over the Contiguous United States. *Water Resources Research*, 55(4). <https://doi.org/10.1029/2018WR022797>
- 65 Crosbie, R. S., Binning, P., & Kalma, J. D. (2005). A time series approach to inferring groundwater recharge using the water table fluctuation method. *Water Resources Research*, 41(1). <https://doi.org/10.1029/2004WR003077>
- Fan, J., Oestergaard, K. T., Guyot, A., & Lockington, D. A. (2014). Estimating groundwater recharge and evapotranspiration from water table fluctuations under three vegetation covers in a coastal sandy aquifer of subtropical Australia. *Journal of Hydrology*, 519(PA). <https://doi.org/10.1016/j.jhydrol.2014.08.039>
- 70 Jasechko, S., Seybold, H., Perrone, D., Fan, Y., & Kirchner, J. W. (2021). Widespread potential loss of streamflow into underlying aquifers across the USA. *Nature*, 591(7850). <https://doi.org/10.1038/s41586-021-03311-x>
- McKay, L., Bondelid, T., Dewald, T., Johnston, J., Moore, R., & Rea, A. (2012). *NHDPlus Version 2: User Guide*.
- Siva Prasad, Y., & Venkateswara Rao, B. (2018). Groundwater recharge estimation studies in a khondalitic terrain of India. *Applied Water Science*, 8(4). <https://doi.org/10.1007/s13201-018-0738-2>

- 75 U.S. Geological Survey. (2019). *3D Elevation Program 1-Meter Resolution Digital Elevation Model*.
- van Genuchten, M. Th. (1980). A Closed-form Equation for Predicting the Hydraulic Conductivity of Unsaturated Soils. *Soil Science Society of America Journal*, 44(5). <https://doi.org/10.2136/sssaj1980.03615995004400050002x>
- 80 Wieczorek, M. E., Jackson, S. E., & Schwarz, G. E. (2019). *Select Attributes for NHDPlus Version 2.1 Reach Catchments and Modified Network Routed Upstream Watersheds for the Conterminous United States USGS data release v. 2.0*. USGS.
- Yin, L., Hu, G., Huang, J., Wen, D., Dong, J., Wang, X., & Li, H. (2011). Groundwater-recharge estimation in the Ordos Plateau, China: Comparison of methods. *Hydrogeology Journal*, 19(8). <https://doi.org/10.1007/s10040-011-0777-3>
- 85 Zhang, Z., Wang, W., Gong, C., Zhao, M., Wang, Z., & Ma, H. (2021). Effects of non-isothermal flow on groundwater recharge in a semi-arid region. *Hydrogeology Journal*, 29(2). <https://doi.org/10.1007/s10040-020-02217-8>

90