

State of Wildfires 2023-24

Authors:

Matthew W. Jones^{1,*}, Douglas I. Kelley^{2,*}, Chantelle A. Burton^{3,*}, Francesca Di Giuseppe^{4,*}, Maria Lucia F. Barbosa^{5,6}, Esther Brambleby¹, Andrew J. Hartley³, Anna Lombardi⁷, Guilherme Mataveli^{8,1}, Joe R. McNorton⁴, Fiona R. Spuler⁹, Jakob B. Wessel^{10,11}, John T. Abatzoglou¹², Liana O. Anderson¹³, Niels Andela¹⁴, Sally Archibald¹⁵, Dolores Armenteras¹⁶, Eleanor Burke³, Rachel Carmenta¹⁷, Emilio Chuvieco¹⁸, Hamish Clarke¹⁹, Stefan H. Doerr²⁰, Paulo M. Fernandes²¹, Louis Giglio²², Douglas S. Hamilton²³, Stijn Hantson²⁴, Sarah Harris²⁵, Piyush Jain²⁶, Crystal A. Kolden²⁷, Tiina Kurvits²⁸, Seppe Lampe²⁹, Sarah Meier³⁰, Stacey New³, Mark Parrington³¹, Morgane M. G. Perron³², Yuquan Qu^{33,34}, Natasha S. Ribeiro³⁵, Bambang H. Saharjo³⁶, Jesus San-Miguel-Ayanz³⁷, Jacquelyn K. Shuman³⁸, Veerachai Tanpipat³⁹, Guido R. van der Werf⁴⁰, Sander Veraverbeke^{33,1}, Gavriil Xanthopoulos⁴¹

Institutions:

1. Tyndall Centre for Climate Change Research, School of Environmental Sciences, University of East Anglia, Norwich Research Park, Norwich, UK, NR4 7TJ
2. Hydro-climate risks, UK Centre for Ecology & Hydrology, Wallingford OX10 8BB, U.K.
3. Hadley Centre, Met Office, Fitzroy Road, Exeter, UK, EX1 3PB
4. Earth System Modelling Section, Forecast Department, European Centre for Medium-range Weather Forecast, Shinfield Park, Reading RG29AX, United Kingdom
5. Department of Remote Sensing, National Institute for Space Research, Avenida dos Astronautas, 1758. Jd. Granja - São José dos Campos - São Paulo, Brazil , 12227-010
6. Natural Sciences Center, Federal University of São Carlos, Rodovia Lauri Simões de Barros, km 12 - SP-189 - Aracaçu, Buri - São Paulo, Brazil, 18290-000
7. Climate Intelligence, Research Department, European Centre for Medium-range Weather Forecast, Shinfield Road, Reading, UK, RG29AX
8. Earth Observation and Geoinformatics Division, National Institute for Space Research, Avenida dos Astronautas, 1758. Jd. Granja - São José dos Campos - São Paulo, Brazil , 12227-010
9. Department of Meteorology, University of Reading, University of Reading, Earley Gate, Whiteknights Rd, Reading RG6 6ET
10. Department of Mathematics and Statistics, University of Exeter, Harrison Building, University of Exeter, North Park Road, Exeter, UK
11. The Alan Turing Institute, British Library, 96 Euston Road, London, UK
12. School of Engineering, University of California, Merced, 5200 N Lake Rd, Merced, CA, 95343, USA
13. Cemaden/MCTI, Estrada Doutor Altino Bondensan, 500 - Distrito de Eugênio de Melo, São José dos Campos - São Paulo, Brazil
14. BeZero Carbon, 25 Christopher Street, London, UK, EC2A 2BS
15. School of Animal Plant and Environmental Sciences, University of Witwatersrand Johannesburg, University Corner, Braamfontein, Johannesburg
16. Landscape Ecology and Ecosystem Modelling Group, Faculty of Sciences, Department of Biology, Universidad Nacional de Colombia, Cra. 30 # 45-03, Bogotá D.C., CP 111321, Colombia
17. Tyndall Centre for Climate Change Research, School of Global Development, University of East Anglia, Norwich Research Park, Norwich, UK, NR4 7TJ
18. Department of Geology, Geography and the Environment, Universidad de Alcalá, Colegios, 2 - 28801 Alcalá de Henares
19. FLARE Wildfire Research, School of Agriculture, Food and Ecosystem Sciences, University of Melbourne, Grattan St, Parkville, Australia, 3010
20. Centre for Wildfire Research, Swansea University , Singleton Park Swansea SA2 8PP Wales, UK
21. ForestWISE—Collaborative Laboratory for Integrated Forest & Fire Management, Centre for the Research and Technology of Agro-Environmental and Biological Sciences, Universidade de Trás-os-Montes e Alto Douro, Quinta de Prados, Vila Real, Portugal, 5000-801
22. Department of Geographical Sciences, University of Maryland, College Park, MD 20742
23. Marine, Earth and Atmospheric Science , North Carolina State University, Raleigh, North Carolina, USA, 27695
24. Program in Earth System Sciences, Faculty of Natural Sciences, Universidad del Rosario, Bogotá, Colombia
25. Fire Risk, Research and Community Preparedness, Country Fire Authority, Burwood East, Victoria, Australia

- 59 26. Northern Forestry Centre, Canadian Forest Service, Natural Resources Canada, 5320 122 St NW,
60 Edmonton, AB T6H 3S5, Canada
61 27. Wildfire Resilience Center, School of Engineering, University of California, Merced, 5200 N Lake Rd, Merced,
62 CA, 95343, USA
63 28. GRID-Arendal, P.O Box 183, N-4802, Arendal, NORWAY
64 29. Department of Water and Climate, Vrije Universiteit Brussel, Pleinlaan 2, 1050 Brussel, Belgium
65 30. Land, Environment, Economics and Policy Institute, Department of Economics, University of Exeter, Rennes
66 Drive, Exeter, United Kingdom, EX4 4ST
67 31. Atmospheric Composition Section, Research Department, European Centre for Medium-range Weather
68 Forecast, Robert-Schuman-Platz 3, 53175 Bonn, Germany
69 32. UMR 6539 CNRS/IRD/Ifremer/LEMAR, Institut Universitaire Européen de la Mer, University of Brest, F-29280
70 Plouzané, France
71 33. Department of Earth Sciences, Faculty of Science, Vrije Universiteit Amsterdam, De Boelelaan 1105, 1081
72 HV Amsterdam, Netherlands
73 34. Institute of Bio- and Geosciences: Agrosphere (IBG-3), Forschungszentrum Jülich, Wilhelm-Johnen-Straße,
74 52428 Jülich, Germany
75 35. Faculty of Agronomy and Forest Engineering, Eduardo Mondlane University, 3453 Avenida Julius Nyerere,
76 Maputo, Mozambique
77 36. Faculty of Forestry, Bogor Agricultural University, Kampus Ipb Darmaga, Bogor
78 37. European Commission Joint Research Center, European Commission, Rue du Champ de Mars 21, 1050
79 Brussels, Belgium
80 38. NASA Ames Research Center, PO Box 1 Moffett Field, CA 94035-1000
81 39. Upper ASEAN Wildland Fire Special Research Unit (WFSRU), Forestry Research Center, Faculty of Forestry,
82 Kasetsart University, 5th Floor, 72nd Anniversary of Faculty Forestry Building
83 40. Wageningen University, Droevendaalsesteeg 3, 6708PB Wageningen
84 41. Forest Fire Laboratory, Institute of Mediterranean Forest Ecosystems, Hellenic Agricultural Organization
85 (DIMITRA), Terma Alkmanos, Ilisia, 11528, Athens, Greece

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87 ★ These authors contributed equally to this work.

88

89 Correspondence to:

90 matthew.w.jones@uea.ac.uk

91 doukel@ceh.ac.uk

92 chantelle.burton@metoffice.gov.uk

93 Francesca.DiGiuseppe@ecmwf.int

94

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96

97 **Abstract**

98

99 Climate change contributes to the increased frequency and intensity of wildfires globally, with
100 significant impacts on society and the environment. However, our understanding of the global
101 distribution of extreme fires remains skewed, primarily influenced by media coverage and
102 regionalised research efforts. This inaugural State of Wildfires report systematically analyses
103 fire activity worldwide, identifying extreme events from the March 2023-February 2024 fire
104 season. We assess the causes, predictability, and attribution of these events to climate
105 change and land use, and forecast future risks under different climate scenarios. During the
106 2023-24 fire season, 3.9 million km² burned globally, slightly below the average of previous
107 seasons, but fire carbon (C) emissions were 16% above average, totalling 2.4 Pg C. Global
108 fire C emissions were increased by record emissions in Canadian boreal forests (over 9 times
109 the average) and reduced by low emissions from African savannahs. Notable events included
110 record-breaking fire extent and emissions in Canada, the largest recorded wildfire in the
111 European Union (Greece), drought-driven fires in western Amazonia and northern parts of
112 South America, and deadly fires in Hawaii (100 deaths) and Chile (131 deaths). Over 232,000
113 people were evacuated in Canada alone, highlighting the severity of human impact. Our
114 analyses revealed that multiple drivers were needed to cause areas of extreme fire activity. In
115 Canada and Greece a combination of high fire weather and an abundance of dry fuels

116 increased the probability of fires, whereas burned area anomalies were weaker in regions with
117 lower fuel loads and higher direct suppression, particularly in Canada. The fire season in
118 Canada was predictable three months in advance based on the fire weather index, whereas
119 events in Greece and Amazonia had shorter predictability horizons. Attribution analyses
120 indicated that the areas burned by fires in Canada, Greece and western Amazonia during the
121 2023-24 fire season were up to 40.1%, 17.7%, and 50% higher due to climate change.
122 Meanwhile, the probability of extreme fire seasons on these magnitudes has increased
123 significantly due to anthropogenic climate change, with a 2.9-3.6-fold increase in likelihood of
124 high fire weather in Canada and a 20.0-28.5-fold increase in Amazonia. By the end of the
125 century, events of similar magnitude to 2023 in Canada are projected to occur 6.3-10.8 times
126 more frequently under a medium-high emission scenario (SSP370). This report represents our
127 first annual effort to catalogue extreme wildfire events, explain their occurrence, and predict
128 future risks. By consolidating state-of-the-art wildfire science and delivering key insights
129 relevant to policymakers, disaster management services, firefighting agencies, and land
130 managers, we aim to enhance society's resilience to wildfires and promote advances in
131 preparedness, mitigation, and adaptation.
132

133 **Short Summary**

134
135 This inaugural State of Wildfires report catalogues extreme fires of the 2023-24 fire season.
136 For key events, we analyse their predictability and drivers and attribute them to climate change
137 and land use. We provide a seasonal outlook and decadal projections. Key anomalies
138 occurred in Canada, Greece, and western Amazonia, with other high-impact events
139 catalogued worldwide. Climate change significantly increased the likelihood of extreme fires,
140 and mitigation is required to lessen future risk.

141 **1 Introduction**

142

143 **1.1 Background**

144

145 The potential for wildfires is growing under climate change, with increases in the frequency
146 and intensity of drought and periods of fire-favourable weather driving reductions in vegetation
147 (fuel) moisture and priming landscapes to burn more regularly, severely, and intensely
148 (Seneviratne et al., 2022; UNEP, 2022a; Jones et al., 2022; Abatzoglou et al., 2019;
149 Cunningham et al., 2024a). Additionally, human activities and land use change can contribute
150 to or exacerbate the risk of extremely large, fast-moving or intense fires, especially in tropical
151 forests where people are the primary cause of ignition and forest degradation (Lapola et al.,
152 2023). Recent years have been marked by a series of extreme wildfire events spanning the
153 globe, with record levels of burned area (BA) occurring in the 2019-2020 Australian "Black
154 Summer" bushfires (Abram et al., 2021) and a series of high-ranking wildfire seasons
155 occurring in quick succession in the western US (2020 and 2021; Higuera & Abatzoglou,
156 2020), Siberia (2020 and 2021; Zheng et al., 2023), Europe (2022; European Commission
157 Joint Research Centre, 2023) and South America (2019, 2020; Kelley et al., 2021; Ferreira
158 Barbosa et al., 2022; Silveira et al., 2020). The 2023-24 fire season was marked by
159 unprecedented fire extent and emissions in Canada, deadly fast-moving fires in Hawaii and
160 Chile, the largest individual wildfires on record in the European Union and Canada, and
161 widespread fires in northwestern South America including parts of Amazonia such as Brazil,
162 Bolivia, Colombia, and Venezuela (Mataveli et al., 2024; Kolden et al., 2024; European
163 Commission EU Science Hub, 2023).
164

165 The prominence of recent extreme wildfires and wildfire seasons notably contrasts with overall
166 trends in the area burned by fires globally. Due mostly to a reduction in the global savannahs
167 tied to landscape fragmentation and changing rainfall patterns, global BA has fallen since the

168 beginning of this century by around one-quarter (Andela et al., 2017; Jones et al., 2022; Chen
169 et al., 2024). Critically, this decline in fire extent masks major shifts in the distribution of fires
170 globally, with regions such as eastern Siberia and the western US and Canada experiencing
171 a more than 40% increase in BA since 2000 (Jones et al., 2022; Zheng et al., 2021) and
172 regions such as southeast Australia also showing significant increases over longer periods
173 (Canadell et al., 2021). Likewise, there have been shifts in the global distribution of BA and
174 fire carbon (C) emissions from non-forests to forests globally and from the tropics to the
175 extratropics (Kelley et al., 2019). Hence, focussing on global aggregated BA extent underplays
176 the scale and magnitude of changes in wildfire activity and impact on regional levels. An
177 increase in forest and peatland burning is particularly concerning due to the rich ecosystem
178 services that these regions provide, including C storage and biodiversity (UNEP, 2022b). The
179 intensification of fire regimes in environments that are less fire-adapted is particularly
180 important because these ecosystems are expected to be least resilient to such changes (Grau-
181 Andrés et al., 2024).

182
183 The extreme wildfire events of recent years have significantly impacted society and
184 ecosystems across the globe (Cunningham et al., 2024a). Since 1990, wildfire disasters have
185 directly killed or injured at least ~18,000 people, a conservative measure based on incomplete
186 records and reporting biased to the global Northern countries (updated from Jones et al., 2022;
187 Centre for Research on the Epidemiology of Disasters, 2024). In 2023, 232,000 people were
188 evacuated due to wildfires in Canada alone (Jain et al., 2024; Kolden et al., 2024). Also since
189 1990, fires are estimated to have caused on the order of 10 million premature deaths globally
190 through degraded air quality (Johnston et al., 2012). Degraded air quality related to fires is
191 experienced most strongly in the tropics (Pai et al., 2022) and often disproportionately affects
192 traditional communities with poor public services or means of protection (Carmenta et al.,
193 2021). Yet, images of North American cities blanketed in smoke during the 2023 fire season
194 highlight the global nature of this problem.

195
196 As anthropogenic emissions of CO₂ remain persistently high, the world's natural C sinks in
197 forests, peatlands and other ecosystems are increasingly pivotal to moderating increases in
198 atmospheric CO₂ concentration (Friedlingstein et al., 2023). Intact forests are often relied upon
199 for delivering national plans for reaching Net Zero (Smith et al., 2023) and offering sites for
200 nature based solutions (NBS). Yet, massive wildfire emissions from boreal forests and soils in
201 Siberia and Canada across the years 2020, 2021, and 2023 amount to over 1 billion tonnes
202 of C, a gross flux comparable in magnitude to annual CO₂ emissions from fossil fuel
203 combustion in India, the EU27 or the USA (Friedlingstein et al., 2023; Zheng et al., 2023).
204 While in a natural fire regime these gross emissions should be recuperated through post-fire
205 recovery, the greater vegetation mortality and loss of ecosystem function associated with more
206 widespread and severe fires can also contribute to shifts in local to regional terrestrial C
207 budgets from sinks to sources (Zheng et al., 2021; Gatti et al., 2021; Nolan et al., 2021a;
208 Phillips et al., 2022; Harrison et al., 2018; Cunningham et al., 2024b). Loss of vegetation during
209 extreme fire seasons can also have wider lasting effects on ecosystems, for instance by
210 reducing the habitat area available to endemic species (Ward et al., 2020).

211
212 Extreme fires can moreover impact the livelihoods of various communities and landowners
213 who depend on intact natural landscapes. For example, the lands and territories of traditional
214 communities and Indigenous Peoples can be degraded and transformed by wildfires, raising
215 climate justice issues (Garnett et al., 2018; Barlow et al., 2018; Lapola et al., 2023). Further,
216 conflating the detrimental impacts of wildfires types has also stigmatised small-scale
217 intergenerational fire use and led to prohibitive fire governance that affects local communities
218 (Carmenta et al., 2021; Barlow et al., 2020).

219
220 Mitigating and adapting to increases in wildfire potential are growing priorities of policymakers
221 and require coordination with many other stakeholders. National and international disaster
222 management centres are seeking to enhance predictive capacity, while fire management

223 agencies are expanding or re-allocating their resources to rapidly suppress fires to avoid them
224 becoming too large, fast or intense. A number of international organisations such as the UN
225 Environment Programme (UNEP, 2022a), the World Bank (2020, 2024), and the Organisation
226 for Economic Co-operation and Development (OECD, 2023) and a range of other inter- or
227 non- governmental organisations are producing reports that consolidate evidence on the
228 changing risk of extreme fires and identify best practices for mitigating their impacts, including
229 through land management and urban/rural planning. Many land managers are developing and
230 implementing approaches such as fuel reduction, a process subject to permit systems issued
231 by regional fire management agencies in some countries (Fernandes and Botelho, 2003;
232 Stephens et al., 2012; Moreira et al., 2020; Chuvieco et al., 2023). Wildfire response agencies
233 are exploring innovative approaches to detecting and responding to fires, and there is rising
234 interest in the prospect of integrated fire management around the world (Food and Agriculture
235 Organization of the United Nations, 2024). Operators of C market projects and forest carbon-
236 conservation initiatives, such as REDD+ are particularly wary of the risks that wildfires present
237 to the permanence of C offsets, which often feature as a key tool in national policies and
238 international initiatives for achieving Net Zero emissions (Barlow et al., 2012; Smith et al.,
239 2023).

240
241 Amidst extreme wildfires and wildfire seasons, stakeholders increasingly turn to scientists for
242 answers to pressing questions that naturally arise. How extreme was this fire event in a
243 historical context? Is climate change to blame? Will we see more wildfires like this in the
244 future? Did land management exacerbate or ameliorate the problem? Can we predict events
245 like this in future to improve early warning? What is the role of climate policy in reducing risk
246 of extreme wildfires in future?

247
248 While observational, statistical, and modelling tools for assessing extreme wildfire drivers and
249 predicting wildfire occurrence are advancing rapidly, their application to studying extreme
250 wildfire seasons or events on timescales relevant to public and political interest remains
251 limited. The State of Wildfires report represents a new initiative to routinely catalogue extreme
252 wildfire events at annual frequency and explain their occurrence and relation to climate
253 change. The report incorporates recent methodological advances in disentangling the drivers
254 of selected extreme wildfire events to fuel dryness, fuel load, and weather, and ignition and
255 suppression factors. By applying these methodological advances in conjunction with models
256 of global change, we quantify the change in likelihood of the past year's events under climate
257 and land use changes. Observable fire metrics (e.g. burned area) are the target variable of
258 our causal inference and attribution work, which thereby advances on more common climate
259 attribution studies that attribute change in fire-favourable meteorological conditions to climate
260 change. Overall, this report capitalises on recent advances in the study of extreme fire events
261 and seasons to provide timely information about shifting fire regimes and their causes. The
262 findings of the report are relevant to policymakers, the media, and the wider public.

263

264 **1.2 Objectives of this Report**

265
266 This inaugural edition of the State of Wildfires report aims to stimulate development of tools
267 for understanding and predicting extreme fires and to deliver actionable information to policy
268 and practice stakeholders and wider society. In this edition we:

- 269
- 270 1. Regionally identify extreme individual wildfires or extreme wildfire seasons of the
271 period March 2003-February 2024, and place them in context of recent trends.
- 272 2. Shortlist a selective number of extremes (extreme individual wildfires or extreme
273 wildfire seasons) with notable impacts on society or the environment, which we term
274 the 'focal events' in this report.
- 275 3. Diagnose the contributions of fuel dryness, fuel load, ignitions and suppression to the
276 occurrence of each focal event.

- 277 4. Assess the capacity of operational predictive systems to predict each focal event.
- 278 5. Attribute each focal event to anthropogenic factors including climate change and land
- 279 use.
- 280 6. Provide an outlook on the probability of extreme events in the coming fire season
- 281 (commencing March 2024).
- 282 7. Project future changes in the probability of each focal event under future climate
- 283 scenarios.

284
285 Key methodologies used to achieve the above objectives are summarised as follows. To
286 address objectives 1 and 2, we build a comprehensive dataset of fire metrics including BA,
287 fire counts, fire C emissions, and individual fire properties (size and rate of growth) for
288 consistent world regions and we quantitatively identify anomalies in these metrics during the
289 past fire season (Giglio et al., 2018; van der Werf et al., 2017; Andela et al., 2019). To address
290 objective 3 and 4, we leverage seasonal to sub-seasonal forecasts of fire weather from the
291 European Centre for Medium-Range Weather Forecasts (ECMWF) and additionally employ
292 two state-of-the-art fire models, *Controls on Fire* (ConFire) and *Probability of Fire* (PoF) (Kelley
293 et al., 2019; McNorton et al., 2024) to pinpoint the causes of the extreme fire events of 2023-
294 24. To address objective 5, we employ projections of fire weather from the Hadley Centre
295 Large Ensemble to attribute change in the fire weather index (FWI) to climate change, and we
296 drive ConFire (Kelley et al., 2019) with outputs from the Joint UK Land Environment Simulator
297 Earth System model (JULES-ES; Mathison et al., 2023) to attribute extreme BA to climate and
298 land use changes (Burton, Lampe, et al., 2023). To address objective 6, we consult predictions
299 of the state of climate modes relevant to fire and present seasonal predictions of FWI from the
300 ECMWF (Di Giuseppe et al., 2024). To address objective 7, we again pair ConFire (Kelley et
301 al., 2019) with JULES-ES (Mathison et al., 2023) to project future changes in BA under several
302 future climate and land use scenarios. comprehensive assessment of past and future extreme
303 wildfire events.

304
305 The State of Wildfires report will be an annually recurring report that can harness and adopt
306 new methodologies brought forward by the scientific community between the annual iterations
307 of the report. Over the coming years and decades, we aim to enhance the tools presented in
308 this report for application in near-real time, thus enhancing our capacity to transfer key insights
309 to decision-makers at the time they most need it.

310 **2 Extreme Wildfire Events of 2023-24**

311

312 **2.1 Methods**

313

314 We catalogued the extreme regional wildfire events or annual fire seasons in the period March
315 2023-February 2024 based on a combination of anomalies in the distribution of several
316 observable fire metrics from Earth observations (**Section 2.1.1**). In this work, the global fire
317 season is defined as occurring in March-February windows oriented around the annual minima
318 of global fire activity in boreal spring (see further details in **Section 2.1.1.2**).

319

320 Due to the diversity of environmental settings in which fires occur and the range of ecological,
321 economic, or societal impacts caused, defining an extreme fire or an extreme fire season
322 remains inherently challenging. To date, extreme fires have commonly been defined by their
323 BA extent, by their feedback on the global climate, and by their socio-economic impact. While
324 an extreme fire event or extreme fire season may be visible as a significant anomaly against
325 historical Earth observations, the scientific community seeks to apply a more comprehensive
326 definition of extreme fire, including its impacts on society and the environment. To catalogue
327 extreme events that were not necessarily visible in Earth observations, regional expert panels
328 were constructed and given responsibility for identifying extreme events of the past fire season
329 (**Section 2.1.3**). The expert panel were given flexibility to identify and catalogue wildfire

330 characteristics or impacts that are considered regionally extreme but are not necessarily
331 captured by Earth observations. Examples of extremes that can be captured by expert
332 assessment (but not by Earth observations) include: suppression challenge; fatalities and
333 structure loss; impacts on human health and wellbeing; impacts on agricultural and other
334 economic sectors; impacts on biodiversity, and; impacts on diverse ecosystem services such
335 as recreation, tourism or other cultural values. Hence, **Section 2.2** identifies a variety of
336 impactful events displaying a broad range of characteristics and impacts that can occur across
337 diverse fire regimes (e.g. Archibald et al., 2009; Cunningham et al. 2024b; Keeley, 2009).
338

339 **2.1.1 Earth Observations of Fire**

340 **2.1.1.1 Input Datasets**

341 We assembled observations of burned area (BA), synonymous with fire extent, for the period
342 March 2002-February 2024 from the National Aeronautics and Space Administration (NASA)
343 product MCD64A1 (collection 6.1). MCD64A1 provides daily BA observations at 500 m spatial
344 resolution with global coverage and is based on retrievals from the Moderate Resolution
345 Imaging Spectroradiometer (MODIS) sensors mounted to the Terra and Aqua satellites (Giglio
346 et al., 2018, 2021).
347

348 We also produced a global record of individual fires for the period March 2002-February 2024
349 by updating the Global Fire Atlas (Andela et al., 2019; Andela and Jones, 2023) through
350 February 2024, driven by the 500m MODIS BA data. The Global Fire Atlas algorithm clusters
351 burned cells into individual fires, tracks their daily progression, and logs attributes such as fire
352 size and mean daily rate of growth. Our updates are provided at Andela and Jones (2024).
353 The Global Fire Atlas is one of several products tracking daily fire progression and identifying
354 individual fires at global scale based on moderate resolution satellite data (Andela et al., 2019;
355 Laurent et al., 2018; Artés et al., 2019). The product uses the MODIS BA product and the
356 smallest unit of disaggregation is 500m and the shortest timestep on which the expansion of
357 a fire can be observed is daily. Given its resolution, Global Fire Atlas is expected to represent
358 the dynamics of large fires better than small, fast-moving fires.
359

360 **2.1.1.2 Uncertainties**

361 In addition, we gathered estimates of fire carbon (C) emissions for the period March 2023-
362 February 2024 from two models driven by Earth observations of active fires or burned area:
363 firstly, the Global Fire Assimilation System (GFAS) product, provided operationally by the
364 Copernicus Atmospheric Services (CAMS) at 0.1 degree spatial resolution and daily temporal
365 resolution (Kaiser et al., 2012; European Centre for Medium-Range Weather Forecasts,
366 2024), and; secondly, the Global Fire Emissions Database (GFED; version 4.1s) product at
367 0.25 degree spatial resolution and daily temporal resolution (van der Werf et al., 2017). GFAS
368 is driven by the fire radiative power (FRP) retrievals in the MODIS active fire product
369 MCD14A1 and biome-level relationships between FRP and biomass consumed based on
370 GFED3 (Kaiser et al., 2012). For the 1997-2016 period, GFED4s is driven by MODIS BA data
371 (MCD64A1 collection 5) supplemented with small fire BA based on MODIS active fire data,
372 and a model for biomass productivity and fuel consumption (van der Werf et al., 2017). For
373 the post-2016 period, emissions are based on active fire detections scaled to emissions using
374 pixel-based scaling factors derived from the 2003-2016 overlapping period.
375

376 We note that the MODIS BA product data used in our analyses of anomalies in BA and
377 individual fire properties (via the Global Fire Atlas) are known to be conservative due to the
378 limitations to detecting small fires (e.g. agricultural fires) based on surface spectral changes
379 at 500m resolution. Recent work has shown that including detections of small active fires
380 increases global BA estimates by 93% (Chen et al., 2023). However, variability and trends in
381
382
383

384 regional BA totals using datasets that include small fires do not differ significantly from the
385 variability and trends present in the MODIS BA product (Chen et al., 2023). Hence, inclusion
386 or exclusion of small fires tends to generate biases in central estimates of BA in one direction
387 or the other, in line with the sensitivity of different sensors to different fire types. Uncertainty
388 in the detection of small fires is larger than in the case of fires detected in the MODIS BA
389 product, due to limited validation (van der Werf et al., 2017). The MODIS BA product with
390 resolution of 500 m is deemed highly suitable for addressing the research questions of this
391 report, which focus on more impactful fires that tend to burn larger areas.

392
393 Uncertainties in fire carbon emissions estimates from GFED4.1s are on the order of $\pm 20\text{--}25\%$
394 at 1 standard deviation for global totals (van der Werf et al., 2017; van der Werf et al., 2010).
395 Uncertainties in GFED4.1s are made up of uncertainties in BA, the amount of biomass
396 consumed per unit BA, and the carbon emitted per unit biomass burned. Revisions to BA input
397 data, discussed above, have tended to influence GFED central estimates of fire C emissions
398 to a greater degree than the uncertainties around central estimates (van der Werf et al., 2017;
399 Chen et al., 2023). Uncertainties in fire carbon emissions estimates from GFAS are on the
400 order of approximately $\pm 25\%$ at 1 standard deviation for global totals. Uncertainties are
401 introduced by missed active fire detections, either below the detection threshold of the MODIS
402 instruments, or not observed during the limited diurnal coverage of Low Earth Orbiting
403 satellites, assumptions made for biome classifications, coefficients used to convert observed
404 thermal anomalies to consumed dry matter, and emission factors used to estimate emitted
405 quantities of carbon and pyrogenic pollutants. Variation in C emissions estimates on the order
406 of approximately 20-60% has been observed in studies comparing multiple emissions
407 products (Wiedinmyer et al., 2023).

408

409 **2.1.1.3 Regional Burned Area, Carbon Emissions and Fire Count Totals**

410

411 We calculated regional totals of BA and C emissions based on a variety of regional layers
412 defined in **Table 1**. The regional layers represent a range of biogeographical boundaries (e.g.
413 biomes), geopolitical boundaries (e.g. countries), and values used in scientific reports (e.g. by
414 the Intergovernmental Panel on Climate Change; IPCC). We calculated monthly totals of BA
415 and fire C emissions for each region by aggregating monthly BA and daily C emissions data,
416 summing the data from the input datasets both spatially and temporally as required. In the
417 case of fire C emissions, we also calculated the mean estimate of fire C emissions from
418 GFED4.1s and GFAS, regionally.

419

420 We adopt a March-February definition of the global fire season (e.g. the latest global fire
421 season spans March 2023-February 2024). Due to an annual lull in the global fire calendar in
422 the boreal spring months, fire season BA totals are least sensitive to the shifts in fire season
423 cutoffs of 1-2 months if the fire season centres on spring (Boschetti and Roy, 2008). This
424 makes the global fire season centred on spring a pragmatic option for the study of interannual
425 variability or trends in fire extent (Boschetti and Roy, 2008). The period March-February is
426 specifically oriented at the end of the austral fire season and before widespread fires have
427 begun in the boreal extratropics. The regions where this global definition of the fire season is
428 most problematic are: northern hemisphere South America, Southeast Asia, and Central
429 America (Giglio et al., 2013).

430

431 In addition, we calculated totals of regional fire counts for each global fire season based on
432 the number of individual fire ignition points present within each region, using ignition point
433 vectors from the Global Fire Atlas. The resolution of the MODIS data supplied to the Global
434 Fire Atlas algorithm is 500 m and hence fires that are smaller in scale are omitted. Regional
435 or national systems may record greater fire counts due to the inclusion of smaller fires.

436

437 **Table 1:** Regional layers to which global Earth observations were disaggregated and used to
 438 define regions with extreme wildfire seasons or extreme individual wildfire attributes. Regional
 439 layers are available from Jones et al. (2024).

Layer	Short Form	Source	Notes
Biomes	NA	Olson et al. (2001)	
Continents	NA	ArcGIS Hub (2024)	
Continental Biomes	NA	See above	Spatial intersect of biomes and continents.
Countries	NA	EU Eurostat (2020)	
UC Davis Global Administrative Areas (GADM) Level 1	GADM-L1	UC Davis (2022)	First sub-national administrative level, such as states of the US or provinces of China. Version 4.1.
Intergovernmental Panel on Climate Change <i>Sixth Assessment Report (AR6) Working Group I (WGI) Reference Regions</i>	IPCC AR6 WGI Regions	Iturbide et al. (2020)	
Global C Project <i>Regional C Cycle Assessment and Processes (RECCAP2) Reference Regions</i>	RECCAP2 Regions	Ciais et al. (2022)	
Global Fire Emissions Database (GFED) Basis Regions	GFED4.1s Regions	van der Werf et al. (2006)	

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441

2.1.2 Identifying Extreme Fire Seasons and Events from Earth Observations

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445

2.1.2.1 Regions with Extreme Wildfire Seasons

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448

Anomalies in BA, fire C emissions, and fire counts in the latest global fire season (March 2023-February 2024) were calculated in several ways:

- 449 (i) as relative anomalies (expressed in %) from the annual mean during all previous March-
 450 February periods since 2002 (2003 for fire C emissions);
 451 (ii) as standardised anomalies (standard deviations) from the annual mean during all previous
 452 March-February periods since 2002 (2003 for C emissions);
 453 (iii) as a rank amongst all March-February periods since 2002 (2003 for fire C emissions),
 454 March 2023-February 2024 inclusive.

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In this report, anomalies in fire C emissions are reported based on the two-model mean estimate from GFED4.1s and GFAS, however anomalies based on the GFED4.1s or GFAS estimates individually are also available via Jones et al. (2024).

460
461
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464

We identified regions in which the latest fire season was potentially classifiable as ‘extreme’ based on the rank of BA, C emissions and fire count amongst all fire seasons. For visualisation purposes, we identified regions in which the latest fire season ranked in the top 5 of all annual fire seasons on record (see **Section 2.2.1**). The BA data for the period March 2002-February 2024 includes 23 fire seasons, while the C emissions data for the period March 2003-February

465 2024 includes 21 fire seasons. Hence, a top-5 ranking translates approximately to a fire
466 season in the upper quartile of those on record.

467

468 We further characterised the onset, peak, and cessation of anomalous monthly BA in March
469 2023-February 2024. First, we identified the month of the event's peak as the maximum
470 difference between monthly BA values in March 2023-February 2024 and the climatological
471 mean monthly values from the prior March-February periods. Thereafter, the event's onset
472 and cessation were defined as the bounds of consecutive months with above-average BA
473 prior to and following the peak but limited to the March 2023-February 2024 period.

474

475 **2.1.2.2 Regions with Extreme Individual Wildfire Attributes**

476

477 We identified regions in which large or fast-moving fires occurred in the latest fire season
478 based on records of individual fires from the Global Fire Atlas. For each region (**Table 1**) and
479 year, we estimated the size of the largest fire, the daily rate of growth of the fire that spread
480 most rapidly, the size of the 95th percentile fire, and the daily rate of growth of the 95th
481 percentile fire. In the Global Fire Atlas, the daily rate of growth for any given fire is determined
482 by calculating the average daily rate of growth at which the fire advanced across all its
483 constituent cells. This method includes cells burned by the head, flank, and backfire and
484 produces lower spread rates than if the calculation were based solely on the cells burned by
485 the head fire.

486

487 Anomalies in each fire attribute were calculated using the same metrics as for BA (see *i-iii*
488 above), and we identified regions in which the latest fire season featured fires with potentially
489 extreme attributes based on the rank of BA and fire C emissions amongst all fire seasons.

490

491 **2.1.3 Identifying Extreme Fire Seasons and Events from Expert Consultation**

492

493 We assembled a panel of regional experts (two from each continent, **Table A1**), to contribute
494 to the identification, description, and characterisation of extreme wildfire seasons or impactful
495 events in the latest fire season. A key role of the expert panel was to catalogue regional events
496 that significantly impacted society or the environment but which may not have been detected
497 by Earth-observing satellites due to issues such as scale, short duration, timing of overpass,
498 and cloud or canopy cover. This includes (but is not limited to) wildfires that impacted society
499 by causing fatalities, evacuations, displacement (e.g. homelessness), direct structure or
500 infrastructure loss or damage, degradation of air or water quality, loss of livelihood, cultural
501 practice or other ways of life, and loss of economic productivity. This definition also includes
502 (but is not limited to) wildfires that impact the environment via disturbance to vulnerable
503 ecosystems, biodiverse areas, or ecosystem services such as C storage. This approach
504 recognises that Earth observations do not provide a complete record of all impactful fires. We
505 do not define ubiquitous quantitative thresholds of impact by any of the measures outlined
506 above, but rather invite in-region experts to identify events that triggered impacts that were
507 sufficient in magnitude to infiltrate public and political discourse. The sources of information
508 available for cataloguing regional events include national/regional fire records, fire service
509 reports, disaster management reports, news reports, and social media. A second key role of
510 our expert panel was to describe and contextualise the impacts of the fire seasons highlighted
511 as extreme by Earth observations or regional assessment (see **Section 2.2.3**).

512

513 The year in review by continent, produced by the expert panel, is presented in **Appendix A**.

514

515 **2.1.4 Context of Recent Extremes: Regional Trends in Burned Area**

516

517 To place recent extremes in the context of fire trends of the past two decades, we update our
518 regional analyses of trends in annual BA from Jones et al. (2022). In contrast to Jones et al.

519 (2022), we present trends that align more closely with global fire seasons, spanning the period
520 March 2002-February 2024 rather than trends over calendar years. We quantified trends using
521 the Theil-Sen slope estimator, which is useful when data may contain outliers or be non-
522 normally distributed making it less sensitive to outliers than a standard least squares
523 regression slope. Changes were calculated by multiplying trends (unit year⁻¹) by the number
524 of fire seasons in the period of coverage for each variable (**Section 2.1.1.2**). Relative changes
525 were calculated as the absolute changes divided by the mean annual BA during the period
526 following Jones et al. (2022) and Andela et al. (2017). The significance of trends was evaluated
527 using the Mann-Kendall test, with a confidence level set at 95%.

528
529 In addition to reporting trends in *total* BA, we also present trends in *forest* BA as these regularly
530 diverge from total BA trends (see **Section 2.2.2**). Forest BA is calculated as described in
531 **Section 2.1.1** but after isolating burned cells in areas with tree cover exceeding 30% in
532 NASA's annual MODIS MOD44B collection 6.0 Continuous Vegetation Field product (250 m)
533 (DiMiceli et al., 2015). The 30% threshold is widely used amongst studies of forest cover
534 change (e.g. Li et al., 2017; Cunningham et al., 2020; Sexton et al., 2016).

535 536 **2.1.5 Shortlisting of Focal Events**

537
538 In later sections of this report, we conducted various analyses to understand the causes and
539 predictability of a selection of extreme wildfire seasons or events during March 2023-February
540 2024 (see **Sections 3-5**). We limited the number of analyses to three globally prominent focal
541 events of the 2023-24 global fire season because the approaches used are not operational
542 and time is required to train and optimise our models regionally.

543
544 In discussion with our expert panel, we prioritised the three events studied in this report by
545 weighing up the anomalies in Earth observations during the latest fire season as well as the
546 impacts that these extremes had on people and the environment. The focal events are notable
547 for their international significance, attracting attention from the media and policymakers both
548 within and beyond their region.

549 550 **2.2 Results**

551 552 **2.2.1 Extreme Fire Seasons and Events of 2023-24**

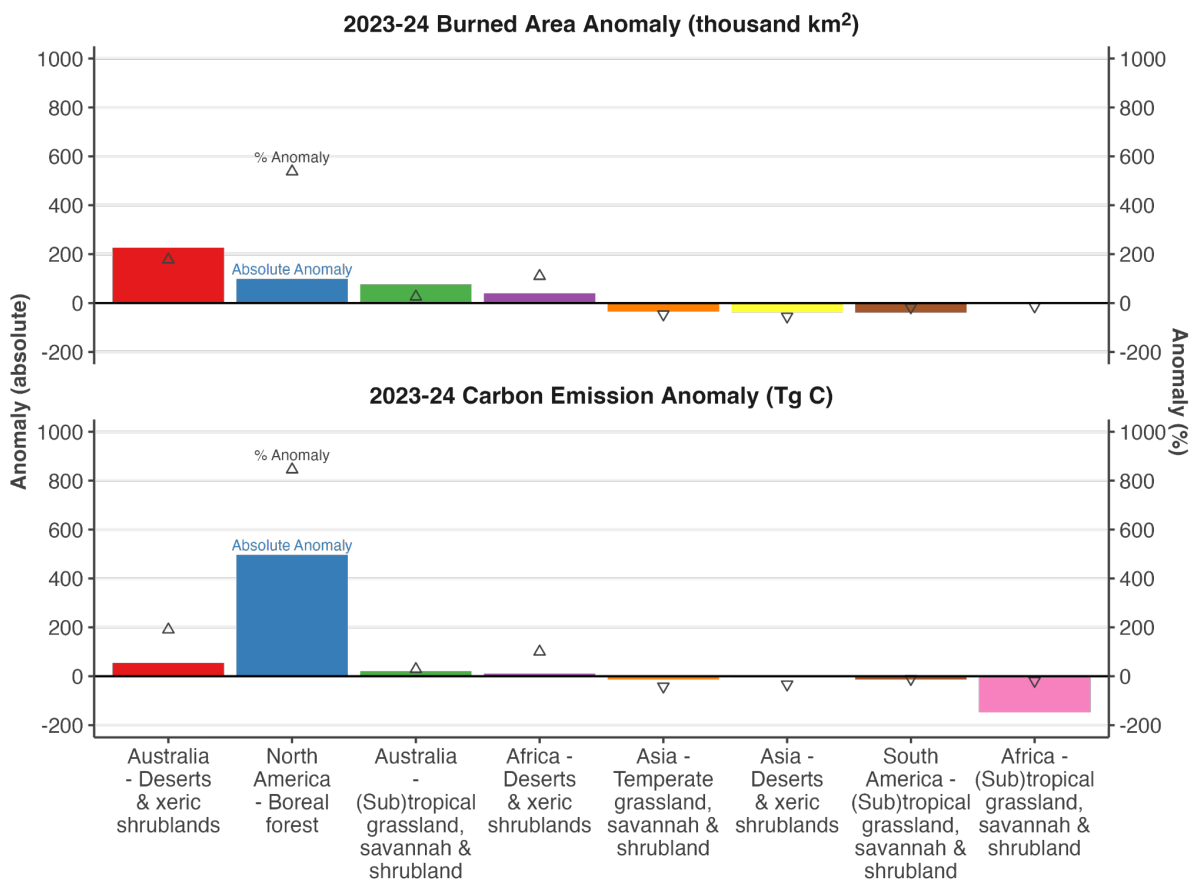
553 554 **2.2.1.1 Extreme Fire Seasons from Earth Observations**

555
556 According to the MODIS BA product, 3.9 million km² burned globally during the 2023-24 global
557 fire season (March 2023-February 2024), slightly below the average of previous fire seasons
558 (4.0 million km²) and overall ranking 12th of all fire seasons since 2002 (Jones et al., 2024).
559 Despite this, fire C emissions were 16% above average at 2.4 Pg C during the 2023-24 global
560 fire season, which ranks 7th amongst all fire seasons since 2003 (based on annual averages
561 of GFED4.1s and GFAS estimates; see methods; Jones et al., 2024).

562
563 Stark regional contrasts in the anomalies in BA, fire C emissions and individual fire properties
564 are visible in the Earth observations on various regional scales (**Figure 1, Figure 2, Figure**
565 **3**). **Figure 1** shows the strongest BA and fire C emissions anomalies of 2023-24 at continental
566 biome scale versus previous fire seasons. BA was around 300 thousand km² (13%) below the
567 average of previous fire seasons in the African grassland, savannah and shrubland biome,
568 which is globally significant because the continental biome contributes 58% towards the global
569 total BA in the average year up to February 2023 (Jones et al., 2024). BA was also around
570 17% below average in the South American grassland, savannah and shrubland biome in 2023-
571 24 and in Asian non-forest biomes. In contrast, BA was 26% above the average of fire seasons
572 since 2002 in the Australian grassland, savannah and shrubland biome (**Figure 1, Figure 2**).

573 Collectively, these three biomes contributed 71% of total BA in the global total BA in the
 574 average fire season between March 2002 and February 2023 and so departures from average
 575 values are particularly impactful on global BA totals.
 576

577 The North American boreal forests experienced a record-breaking fire season, with BA
 578 reaching six times the average since 2002 and fire C emissions reaching over nine times the
 579 average since 2003 (**Figure 1**; Jones et al., 2024). This strong regional signal primarily
 580 explains the above-average global C emissions total of 2023-24, with the high rates of fire
 581 emissions per unit area in boreal forests aggregating to override the reduced emissions totals
 582 in African and South American savannahs. Record levels of fire C emissions were seen also
 583 across the global pan-boreal forest biome, with fire C emissions surpassing the pan-boreal
 584 record set in 2021 by more than 60%. This is despite a below-average fire season for BA and
 585 fire C emissions in boreal Asia during 2023-24, in contrast to the 2021-22 fire season when
 586 there was a synchronous peak in BA in both the Eurasian and North American boreal regions
 587 (Zheng et al., 2021). According to the Global Fire Atlas, new records for individual fire size
 588 and rate of spread were also set in the North American boreal forests during 2023-24, while
 589 95th percentile fire size and rate of growth in 2023-24 were in the top 2 and 3 years on record
 590 since 2002, respectively (Jones et al., 2024). Overall, the Canadian boreal forests contributed
 591 24% towards total fire C emissions in the 2023-24 fire season, up from 3% in an average fire
 592 season since 2003.
 593



594 **Figure 1:** Anomalies in BA and C (C) emissions for selected continental biomes in the 2023-
 595 24 global fire season (March 2023-February 2024), versus the average of prior fire seasons
 596 since 2002. The selected regions all contribute at least 0.1% towards global mean annual BA
 597 and experienced BA anomalies of over $\pm 30,000$ km² in the 2023-24 global fire season. Relative
 598 changes (%) are also marked by triangular symbols and can be read on the same scale as
 599 the absolute values.
 600
 601

602
603 Anomalies in the African (sub-)tropical grasslands, savannahs and shrublands strongly drive
604 inter-annual variability in global fire C emissions because this biome contributes on average
605 40% towards total global fire C emissions (Jones et al., 2024). If fire C emissions from African
606 (sub-)tropical grasslands, savannahs and shrublands had been around average fire season
607 in 2023-24, then global fire C emissions would have been the greatest of any fire season on
608 record since 2003.

609
610 Elsewhere at the biome scale, BA extent was in the top three years on record in the South
611 American broadleaf and mixed forests, the African xeric shrublands, and Australian xeric
612 shrublands, and the Australian (sub-)tropical grasslands, savannahs and shrublands (**Figure**
613 **1**). In contrast, BA or fire C emissions were the lowest on record in the European temperate
614 broadleaf and mixed forests and Asian xeric shrublands, and in the bottom three years on
615 record in the African savannahs, Asian montane grasslands and shrublands, and European
616 tropical grasslands and shrublands.

617
618 On national levels, the most prominent global anomaly of the 2023-24 fire season occurred in
619 Canada where BA reached six times the average of previous fire seasons and fire C emissions
620 reached nine times the average of previous fire seasons. Across the Canadian provinces and
621 territories, the highest BA or fire C emissions on record were observed in Northwest
622 Territories, British Columbia, Alberta, and Quebec while Yukon, New Brunswick, and Ontario
623 also experienced high-ranking years (**Figure 2, Figure 3**). The positive BA anomalies in
624 Canada were visible in the MODIS BA dataset from as early as April 2023 in most provinces
625 and persisted throughout summer through to October and even through to December 2023
626 and January 2024 in British Columbia and Alberta (**Figure S2**). Peak anomalies were
627 observed in Eastern Canada in June 2023, arriving later in western Canada (August-
628 September). Data on individual fire characteristics from the Global Fire Atlas further reveals
629 new record fire counts in many Canadian provinces, and high-ranking anomalies in fire count
630 and daily rate of growth across Canada, as well as new records for fire size and rate of spread
631 in provinces of both eastern and western Canada (**Figure 4**; Jones et al., 2024). **Appendix A**
632 discusses the unprecedented Canadian fire season of 2023-24 in greater detail, including its
633 impacts and regional context.

634
635 A second prominent regional feature of the 2023-24 global fire season, visible in Earth
636 observations, is a cluster of administrative regions with positive BA and C emissions
637 anomalies in the north and west of tropical South America (**Figure 2, Figure 3**). Bolivia,
638 Guyana, Suriname and French Guiana, Honduras, Nicaragua and Belize all experienced a
639 high-ranking fire season at national level in 2023-24. In addition, BA or fire emissions were
640 ranked in the top three years in western parts of Amazonia including in Amazonas state of
641 Brazil, the Loreto department of Peru, and the La Paz and Beni departments of Bolivia.
642 Anomalies in the western Amazon spanned June 2023-January 2024, peaking in August-
643 October 2023. In the north of South America, high-ranking fire seasons were seen in
644 Venezuela, various subdivisions of Guyana, Suriname, and French Guiana, and in Amapá
645 State in Brazil. The anomalies in northern South America spanned May 2023-February 2024,
646 peaking in November 2023-February 2024 (**Figure S2**). The Global Fire Atlas data suggest
647 that South American anomalies in BA during the 2023-24 fire season were principally driven
648 by a large number of fires, whereas anomalies in fire size or rate of growth were uncommon
649 in most of South America (**Figure 4**). **Appendix A** discusses the 2023-24 fire season of
650 tropical South America and its impacts and regional context in greater detail.

651
652 Several parts of South and Southeast Asia experienced high-ranking anomalies in BA or fire
653 C emissions during the 2023-24 fire season, including various neighbouring administrative
654 zones of Lao People's Democratic Republic (PDR), Thailand and Vietnam. The temporal peak
655 of these anomalies was broadly in March-May 2023. Data on individual fire characteristics
656 indicates that high-ranking fire counts, rather than anomalies in fire size, were the primary

657 driver of the regional BA anomalies (**Figure 4**). **Appendix A** discusses these anomalies and
658 their impacts in greater detail.

659

660 The anomalies observed in xeric biomes of Oceania are also apparent as high-ranking BA or
661 C emissions in the 2023-24 fire season in western parts of Australia, particularly in Western
662 Australia and the Northern Territory (**Figure 2, Figure 3**). Fires tended to affect more remote
663 areas and so the impacts on society were muted in comparison to the Black Summer events
664 affecting southeast Australia in 2019-20 (Abram et al., 2021); however, **Appendix A** discusses
665 some notable exceptions.

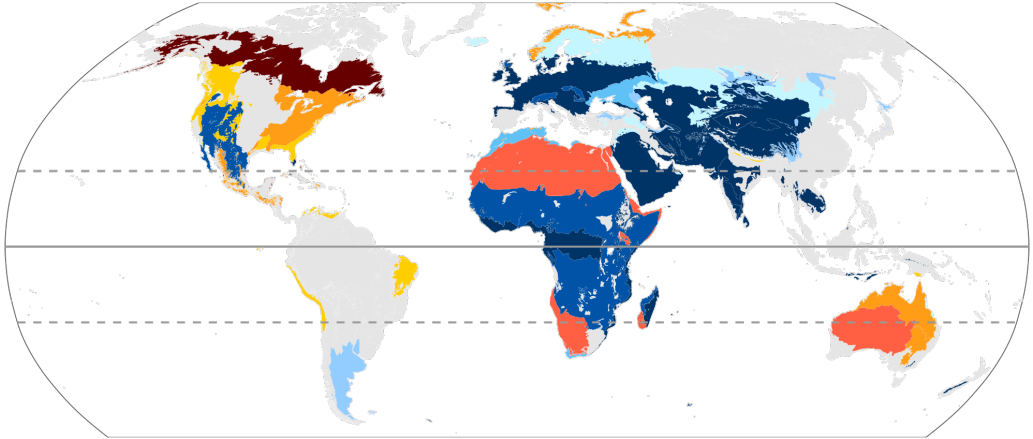
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667 Other regional pockets of high-ranking BA anomalies or C emissions anomalies were
668 observed in various dry zones of Africa and the Middle East, including the Sahel, Northern
669 Africa and the Horn of Africa, Southern Africa (specifically South Africa and Botswana where
670 three high rainfall years have resulted in grass fuel accumulation), parts of Iran, Iraq, parts of
671 the Levant region, and parts of the Arabian Peninsula (**Figure 2, Figure 3**). Although various
672 aspects of the fire season ranked highly in these regions, they are also fuel-limited with a
673 generally a low baseline for BA and fire C emissions and the wildfire season. Nonetheless,
674 regionally impactful wildfires were reported for Algeria, Tunisia and Morocco as well as coastal
675 regions of South Africa and in Pakistan and are discussed further in **Appendix A** and
676 **Appendix A**.

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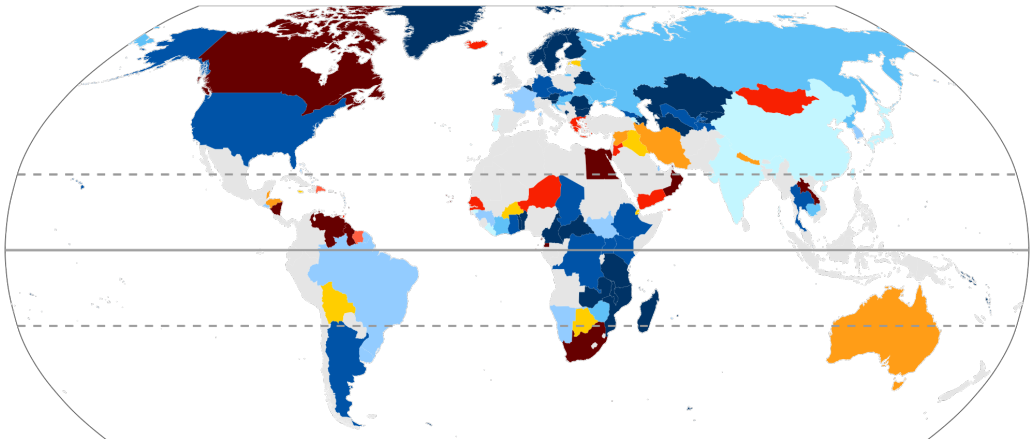
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Continental Biomes



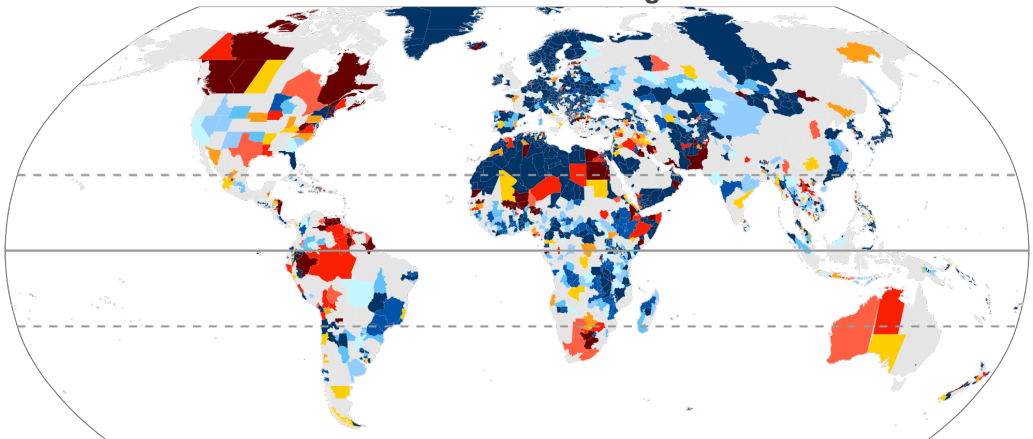
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Countries

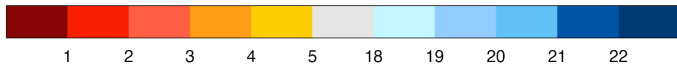


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Level 1 Administrative Regions



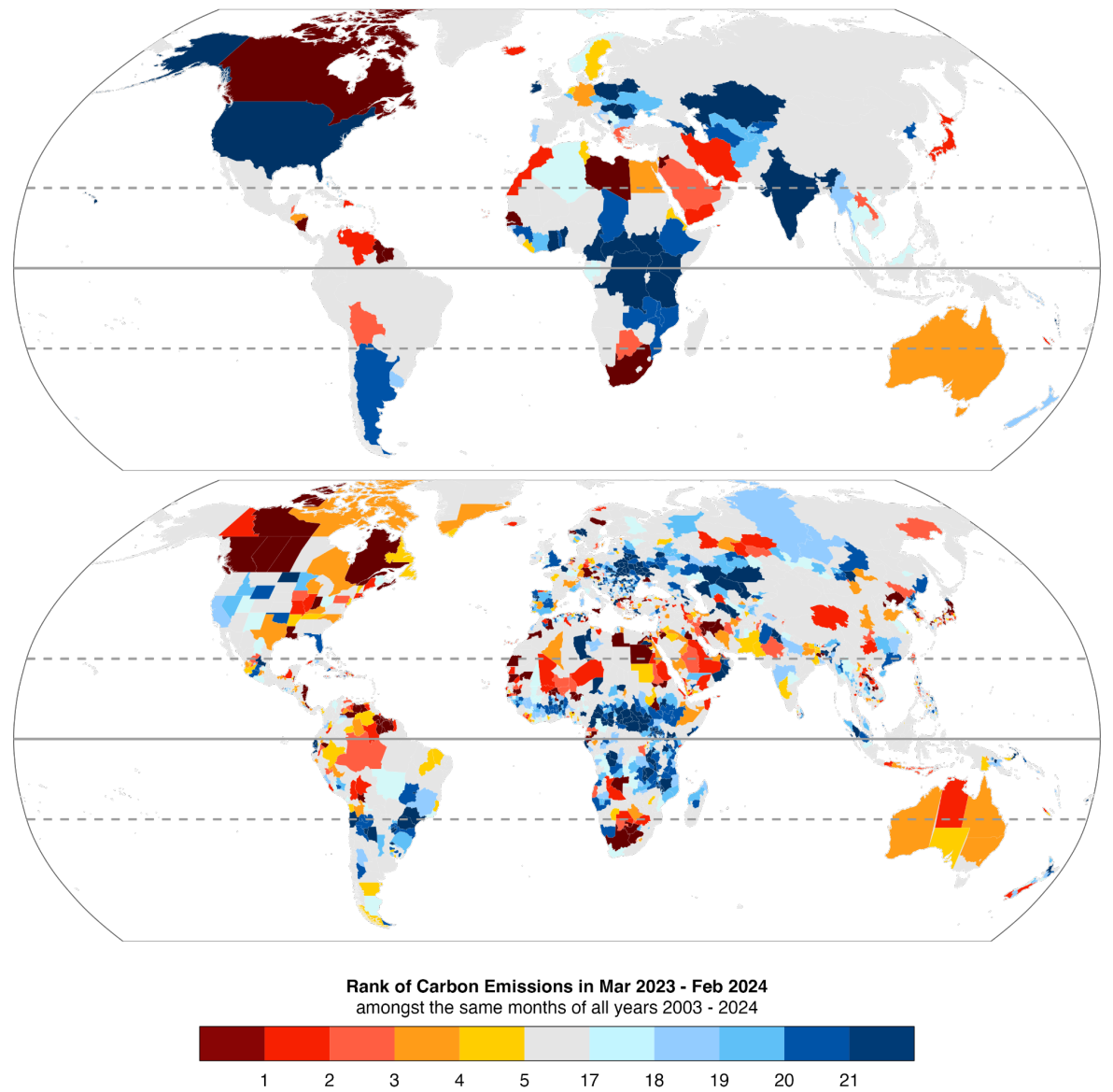
Rank of Burned Area in Mar 2023 - Feb 2024
amongst the same months of all years 2002 - 2024



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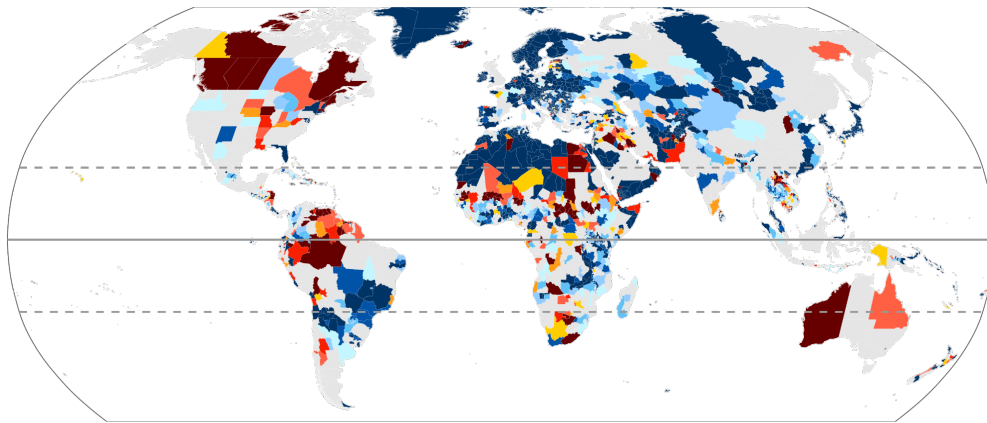
Figure 2: Ranks of BA during March 2023-February 2024 versus previous March-February periods (n = 23 global fire seasons) for three regional layers: **(top panel)** continental biomes; **(middle panel)** countries, and; **(bottom panel)** states or provinces. Results for regions with high-ranking (top 5 years) or low-ranking (bottom 5 years) events are highlighted. The timing of BA anomalies is shown in **Supplementary Figure 2**.

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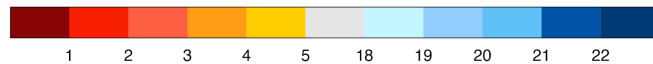


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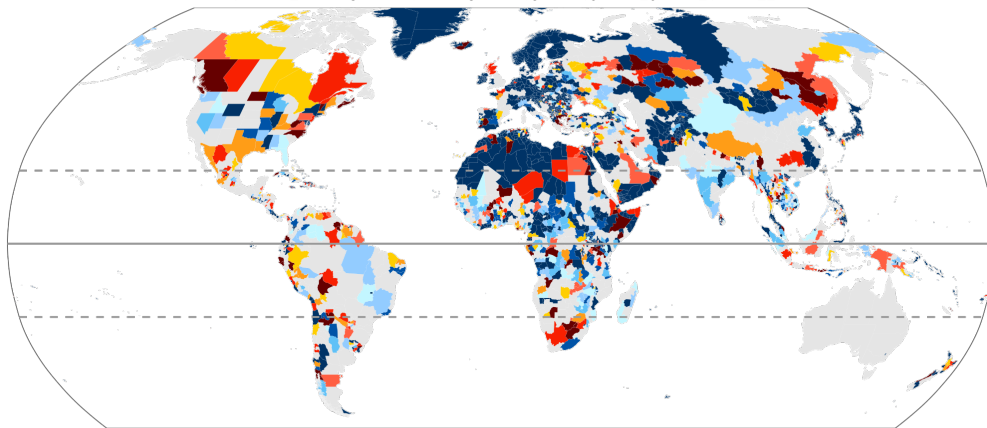
Figure 3: Rank of fire C emissions during March 2023-February 2024 versus all March-January periods since 2003 ($n = 21$ global fire seasons), at the scales of **(top panel)** countries and **(bottom panel)** level 1 administrative regions. We consider C emissions estimates from two products (GFAS and GFED), first calculating the mean emissions value from the two products, then ranking the values.



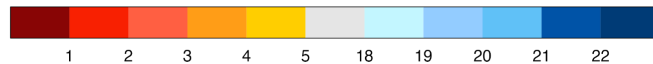
Rank of Number of Fires in Mar 2023 - Feb 2024
amongst the same months of all years 2002 - 2024



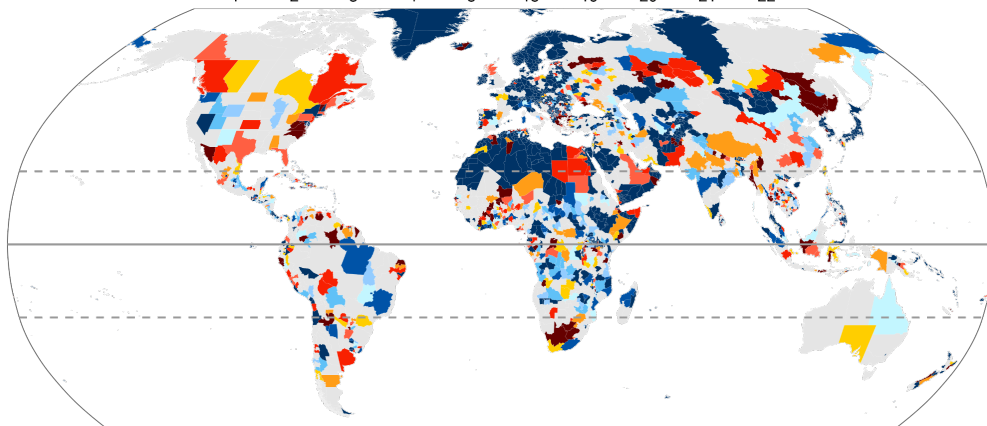
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Rank of 95th Percentile Fire Size in Mar 2023 - Feb 2024
amongst the same months of all years 2002 - 2024



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Rank of 95th Percentile Rate of Growth in Mar 2023 - Feb 2024
amongst the same months of all years 2002 - 2024



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Figure 4: Ranks of **(top panel)** fire count, **(middle panel)** 95th percentile fire size, and **(bottom panel)** 95th percentile daily rate of growth during March 2023-February 2024 versus all March-February periods since 2002, at the scale of states or provinces (GADM administrative level 1 regions).

707 **2.2.1.2 Extreme Individual Fires from Earth observations**

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To support our analyses of anomalies in individual fire properties and provide insights into the limitations and uncertainties inherent in global-scale analysis of individual fires, we provide a brief assessment of the skills with which the Global Fire Atlas represents some of the most impactful individual fires of 2023-24. The Global Fire Atlas represents some of the most impactful individual fire events in 2023-24 with varying skill (**Table 2; Figure S3, Figure S4, Figure S5**). For example, the Evros fire that occurred in the decentralised administration of Macedonia and Thrace, Greece, in late August was captured reasonably well. The Global Fire Atlas identifies two fires that ignited on 19th August and merged to form one contiguous burned unit with an area of approximately 900 km². Alignment of the fire's timing, size and perimeter with high-resolution satellite imagery (**Figure S3**) and detailed reports (Xanthopoulos et al., 2024) suggest an overall reliable representation of this particular event by the Global Fire Atlas. The impacts of this fire are discussed in detail in **Appendix A**.

A deadly fire near Valparaíso, Chile, is also captured with reasonable skill in the Global Fire Atlas (**Figure S4**). Around 90 km² was burned, with the fire skirting the city of Placilla de Peñuelas and encroaching upon Viña del Mar and Quilpué (**Appendix A**). The timing, extent and perimeter of the fire as recorded by the Global Fire Atlas compares well with those reported by other sources (**Table 2**).

Among the largest fires to occur in Canada during 2023-24 happened near the La Grande Reservoir in Quebec, Canada. According to both the Global Fire Atlas and a separate NASA fire tracking product based on the Visible Infrared Imaging Radiometer Suite (VIIRS) sensor, the fire's extent was around 11 thousand km², whereas the National BA Composite (NBAC; Skakun et al. 2022) shows a similar extent of 10 thousand km². The timing of the fire also showed high correspondence across the products.

The Lahaina wildfire in Maui, Hawaii, is an example of an event that was captured poorly by the GFA. Issues relating to the small scale of this fire relative to the resolution of the MODIS BA data are evident in **Figure S4**. As the MODIS BA algorithm is focussed on the detection of wildland fire, its effectiveness in tracking fires at the wildland-urban interface is limited. In this case, burned areas were not detected in cells in urban areas or at the wildland-urban interface, and hence the size of the fire was under-estimated significantly (**Table 2**). The timing of the fire on vegetation land adjacent Lahaina was compatible with reference reports.

Another example of the challenges of defining individual fire extent and applying global algorithms to do so comes from Western Australia (**Figure S5**). Two fires recorded by the Department of Fire and Emergency Services, Western Australia (the Great Sandy Desert and Anna Plains fires) totalled 57 thousand km² in extent. In the Global Fire Atlas, the burned cells detected by MODIS were instead dissected into 53 separate fires with the largest unit burning 560 km². The date ranges were also rather different with the first record of fires logged in agency data one month later than in the Global Fire Atlas and the final record logged one month earlier.

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Table 2: Representation of selected individual fire events in the MODIS BA product (Giglio et al., 2018) and Global Fire Atlas (Andela et al., 2019).

Event	Global Fire Atlas Fire Size (km ²)	GFA Dates	Reported Area (km ²) (Reference)	Dates (Reference)	Reference	Comment
Alexandroupolis Wildfire, Evros, Greece	892	19/08/2023 to 30/08/2023	930	19/08/2023 to 31/08/2023	Xanthopoulos et al. (2024)	Good characterisation of the event, with two fires merging in the date range and ultimate fire size comparable to reference.
Fire near La Grande Reservoir, Quebec, Canada	10,725	29/05/2023 to 23/07/2023	11,400 (VIIRS) 9,694 (NBAC)	01/06/2023 to 23/07/2023	NASA Earth Observatory (2023); VIIRS; Jain et al. (2024; NBAC)	Reasonable characterisation of the fire's extent and timing.
Valparaíso Wildfire, Chile	91.54	31/01/2024 to 10/02/2024	85	02/02/2024 to 05/02/2024	NASA Earth Observatory (2024); Copernicus Emergency Management Service (2023a)	Good characterisation of the scale of the event and its perimeters at various wildland-urban interfaces, versus reference data.
Lahaina Wildfire, Maui, Hawaii	1.50	08/08/2023 to 12/08/2023	8.49	08/08/2023 to 09/08/2023	Fire Safety Research Institute (2024)	MODIS data has coarse spatial resolution relative to scale of event. Spread into urban areas not captured.
Western Australia (Great Sandy Desert and Anna Plains fires)	45,544	31/08/2023 to 28/11/2023	56,561	27/09/2023 to 01/11/2023	Department of Fire and Emergency Services, Western Australia (Shapefile for the Great Sandy Desert and Anna Plains fires; Agnes Kristina, pers. comm.)	Global Fire Atlas splits this event into 53 fires; we report their total combined area. Largest individual fire area in Global Fire Atlas was 760 km ² (ignited 06/09/2023). Great Sandy Desert and Anna Plains fires merged on 25th October 2023.

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2.2.2 Context of Recent Extremes: Regional Trends in Burned Area

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The anomalies of 2023-24 occur against a backdrop of trends in BA this century that point towards shifts in fire regime. **Figure 5** shows significant trends in BA and forest BA across the fire seasons in the period March 2002-February 2024 derived from MODIS BA data. While many world regions are experiencing declines in total BA, increases in forest BA are far more prevalent than declines at the scale of continental biomes, countries, and administrative regions.

Northern hemisphere extratropical biomes in North America and Asia show a clear signal towards increased forest BA since 2002, which are also visible on national and provincial scales in Canada, the US and Russia and on provincial scales in various states of western

768 and eastern Canada, the western US, and Siberia. These trends occasionally propagate to
769 trends in total BA, for example in western and northern Canada and in the Sakha Republic
770 (eastern Siberia). The large 2023-24 anomalies in BA in Canada align with the doubling of
771 forest BA in Canada across fire seasons since 2002 (a significant trend, $p < 0.05$) and a 23%
772 increase in total BA in Canada (marginally significant at $p < 0.1$). Three Canadian provinces
773 showed significant increases in both total and forest BA this century: British Columbia (+35-
774 42%); Northwest Territories (+55-68%), and; Yukon (+60-135%). No Canadian provinces
775 experienced a significant decline in forest BA or total BA. More widely, there was a 58%
776 increase in forest BA in the North American boreal forest biome since 2002, and a 134%
777 increase across the pan-boreal forest biome of North America and Eurasia. The succession
778 of events affecting boreal forests in Canada in 2023, Siberia in 2020, and both North America
779 and Asia during 2021 appear to be part of a continued trend towards rising fire extent in the
780 high latitudes this century.

781
782 Elsewhere in the southern hemisphere extratropics, significant rises in forest BA are seen in
783 Chile since 2002 (+87%), including in the central regions of Araucanía, Bio bío, Maule, Ñuble,
784 and Valparaíso ranging from 35 to 109%. Extreme fires in Valparaíso during 2023-24 and in
785 Araucanía, Biobío and Ñuble in the 2022-23 fire season Maule, Ñuble, Bio bío, Araucanía
786 (**Appendix A**) are also consistent with a longer-term rise BA in central Chile (**Figure 5**).
787 Increases in BA are not generally significant in fire-prone parts of the southern hemisphere
788 extratropics, such as Africa or Australia.

789
790 In the tropics, trends in total and forest BA show a variety of patterns. While total BA has
791 reduced across much of the savannah-occupied regions of South America, Africa and northern
792 Australia, trends in forest BA (>30% tree cover) are far more varied (**Figure 5**). Hence, fires
793 in woody tropical vegetation show a less consistent global trend. In addition, exceptions to the
794 general decline in total BA across the tropics are seen in the Brazilian state of Amazonas, the
795 Congo basin, and across much of India (**Figure 5**). The trend in Amazonas, among the most
796 pristine parts of Amazonia, contrasts with other states of Brazil such as Mato Grosso and
797 Pará, where deforestation rates and deforestation-related fires have fallen since their peak
798 during the early 2000s (Silva Junior, 2020). The anomalous fire activity and C emissions in
799 Amazonas state during the 2023-24 fire season (but not other states of Brazil) thus appear to
800 be consistent with the emerging pattern of increased fire in the region. Meanwhile, the 2023-
801 24 anomalies in BA and other fire properties in the Bolivian, Peruvian, Colombian and
802 Venezuelan parts of Amazonia (**Appendix A**) typically occurred against a backdrop of reduced
803 BA or no significant trend in recent decades (**Figure 5**).

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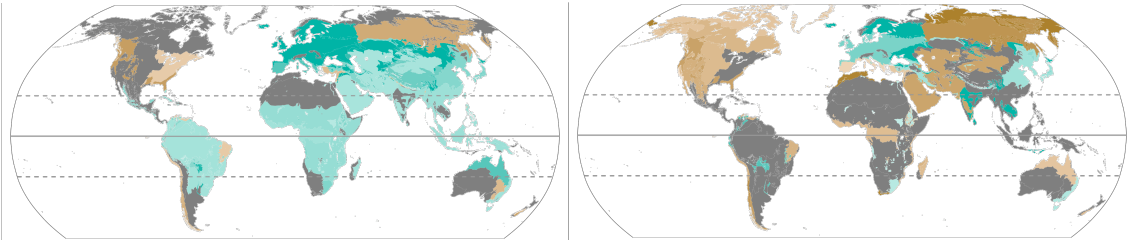
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TOTAL BA

FOREST BURNED AREA

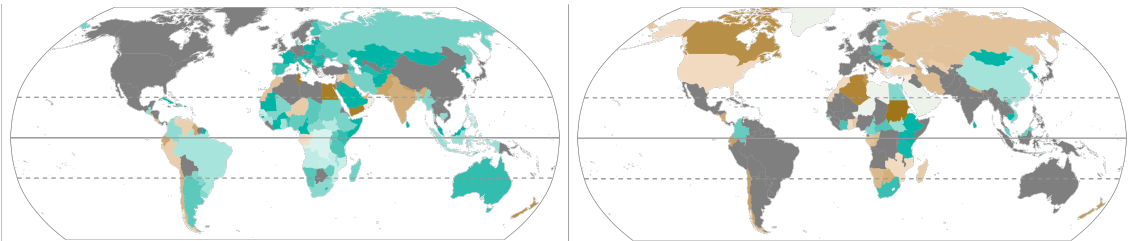
(Forest > 30% tree cover)

Continental Biomes



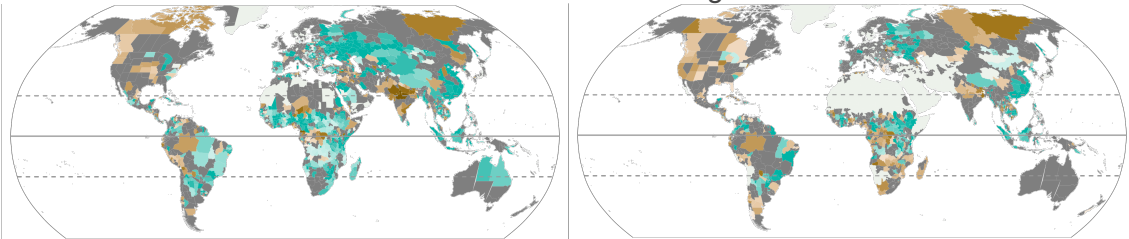
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Countries



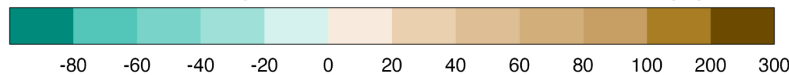
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Level 1 Administrative Regions



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Relative Change in Annual Burned Area, 2002-2024 (%)



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821

Figure 5: Relative changes (%) in **(left panels)** total annual BA and **(right panels)** forest BA across March-February fire seasons during 2002-2024 for three regional layers: **(top panels)** continental biomes; **(middle panels)** countries, and; **(bottom panels)** level 1 administrative regions (e.g. states or provinces). Forest BA considers only areas with tree cover over 30% at the native (500 m) resolution of the BA observations. Relative changes are calculated as the trend in BA across fire seasons March 2002-February 2022 through March 2023-February 2024 multiplied by the number of years in the time series and divided by the mean annual BA during the period. Trends in BA are derived using the Theil-Sen slope estimator. Only significant trends in BA are shown (dark grey fill signifies no significant trend).

831

2.2.3 Focal Events of this Report

832

2.2.3.1 Canada

833

834

In this year's report, the extreme wildfire season in Canada is selected as one of our focal events. It emerges as a major event of global relevance for the following reasons (see **section 2.2.1** and the results of the expert consultation presented in **Appendix A**):

838

839

840

- **Record-Breaking Burned Area:** The North American boreal forests, particularly in Canada, experienced an unprecedented fire season. The BA reached six times the average since 2001.

841

842

- 843 ● **High C Emissions:** Fire C emissions in Canada were over nine times the average
844 since 2003, contributing significantly to global C emission totals for the year. Canadian
845 boreal forests contributed 24% towards the total above-average global fire C emissions
846 in 2023-24, up from 3% in an average year.
- 847 ● **Early and Persistent Fires:** Positive BA anomalies were visible from April 2023
848 (**Figure S9**) and persisted through to October, with some regions experiencing fires
849 until January 2024. The fire season lasted nearly a month longer than normal, with the
850 largest one-day BA total ever recorded in Canada occurring on 22nd September 2023.
- 851 ● **Regional Anomalies:** Peak fire anomalies were observed in Eastern Canada in June
852 2023 and later in Western Canada (August-September), indicating widespread and
853 prolonged fire activity across the country.
- 854 ● **Record Fire Size and Spread:** New records for individual fire size and rate of spread
855 were set, with many provinces experiencing high-ranking anomalies in fire count and
856 daily growth rates.
- 857 ● **Extensive Impact Across Provinces:** The highest BA or fire C emissions on record
858 were observed in Northwest Territories, British Columbia, Alberta, and Quebec, with
859 other provinces like Yukon, New Brunswick, and Ontario also experiencing significant
860 fire activity.
- 861 ● **Air Quality Impact:** Smoke from these fires led to severe air quality issues, affecting
862 major cities in North America, including New York, which experienced its worst air
863 quality in half a century.
- 864 ● **Firefighting Challenges:** Canada was at its highest National Preparedness Level for
865 an unprecedented 120 continuous days, indicating the significant resource sharing and
866 international assistance required to manage the fires.
- 867 ● **Human and Economic Toll:** Over 232,000 people were evacuated across various
868 regions, and despite the extreme fire activity, no civilian deaths were directly attributed
869 to the fires, showcasing the effective, albeit strained, emergency response efforts.

870
871 To assess the causes of specific regional BA anomalies, four anomalous BA regions/month
872 combinations were chosen across Canada: Western Taiga Shield and Taiga Boreal Plains for
873 May and June (includes Alberta and British Columbia Boreal plains, and the Mountain
874 Cordillera); and Eastern Taiga Shield in Quebec for June and July. **Figure 6** maps the
875 magnitude of anomalies in these regions and months. Though note the size and long period
876 this protracted event means that even these regions/month do not capture all the anomalous
877 BA over Canada in 2023 (**Figure S9**).

878 879 **2.2.3.2 Greece**

880
881 In this year's report, the extreme wildfire season in Greece is selected as one of our focal
882 events. It emerges as a major event of global relevance for the following reasons (see **Section**
883 **2.2.1** and the results of expert consultation provided in **Appendix A**):

- 884
885 ● **Second-Highest BA on Record:** Greece experienced its second worst fire season in
886 terms of total area burned, with 1,727 km² affected, despite recent efforts to strengthen
887 firefighting mechanisms. The 2023 fire season was notably more severe than typical
888 years, with the total BA significantly exceeding the country's historical averages and
889 recent challenging fire seasons.
- 890 ● **Multiple Large Fires:** From mid-July to late August, Greece faced numerous large
891 fires that overwhelmed firefighting capabilities. Key fires included those on the island
892 of Rhodes, which burned 207 km², and the massive Evros fire, which reached 938 km².
- 893 ● **Evros Fire Disaster:** The Evros fire became the largest on record in recent European
894 history, significantly impacting both forested and agricultural areas. It also led to the
895 tragic deaths of 19 immigrants who were trapped by the flames.
- 896 ● **Urban and Infrastructural Impact:** Fires near populated areas necessitated large-
897 scale evacuations, including 20,000 tourists on Rhodes and multiple settlements

- 898 around Mount Parnis in Attica. The Evros fire also caused a powerful explosion at an
899 Air Force base, resulting in damage to the town of Nea Aghialos.
- 900 ● **Significant Evacuations:** Numerous evacuations took place, highlighting the severity
901 of the situation. These included evacuations in Alexandroupolis and its surrounding
902 villages due to the Evros fire.
 - 903 ● **Economic and Environmental Damage:** The fires caused extensive damage to
904 properties, infrastructure, and natural reserves, with significant impacts on biodiversity
905 and local economies.
 - 906 ● **Firefighting Challenges:** The simultaneous spread of multiple fires stretched
907 firefighting resources to their limits, with a notable focus on evacuations rather than
908 fire suppression in some instances.

909
910 Abnormally high burned areas were reported around Alexandroupolis in August and extended
911 further west across the administrative region of Macedonia and Thrace. Anomalies were also
912 present in central Greece and around Athens in July and August (**Figure S10**). **Figure 6** maps
913 the magnitude of the anomalies for August.

914 **2.2.3.3 Western Amazonia**

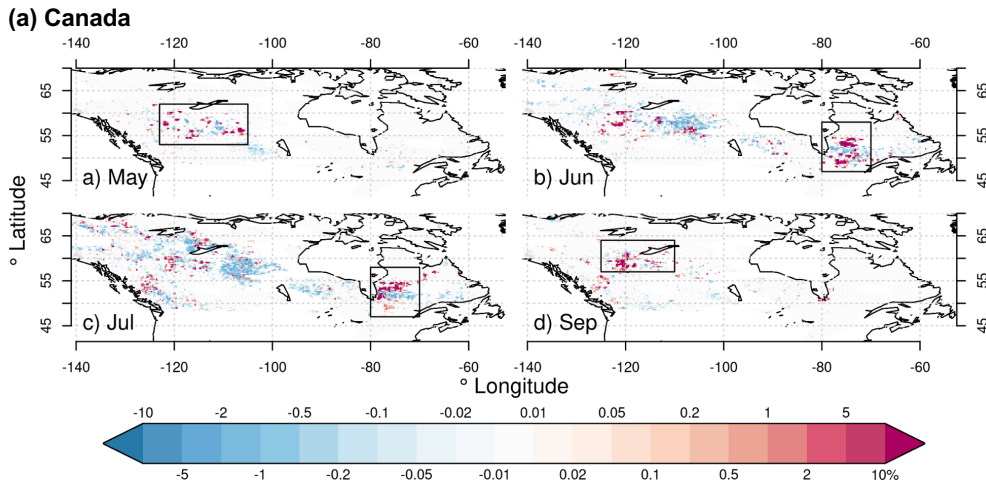
915
916 Our final focal event of 2023-24 is a box drawn in western Amazonia with bounding
917 coordinates 2.25° N, -56.00° E, -9.75° S, -77.75° W. It includes Amazonas (Brazil), Loreto
918 (Peru), and La Paz and Beni (Bolivia) where peak fire anomalies occurred simultaneously. It
919 emerges as a major event of global relevance for the following reasons (see **section 2.2.1**
920 and the results of expert consultation provided in **Appendix A**):

- 921 ● **Record-Setting Fire Activity:** The 2023 fire season in Western Amazonia saw
922 unprecedented fire counts, with new records set across Amazonas state in Brazil,
923 Loreto department in Peru, and La Paz and Beni departments in Bolivia.
- 924 ● **Severe Air Quality Degradation:** Smoke from widespread fires led to significantly
925 degraded air quality across the region, impacting millions of people and posing serious
926 public health risks.
- 927 ● **Broad Socio-Economic and Health Impacts:** The fires caused extensive socio-
928 economic disruptions, including health issues from poor air quality, legal actions for
929 inadequate fire prevention, and impacts on livelihoods, particularly for Indigenous and
930 traditional communities.
- 931 ● **Widespread Environmental Degradation:** The fires contributed to significant C
932 emissions and environmental degradation, affecting forest ecosystems and increasing
933 the region's vulnerability to future climatic extremes. Western Amazonia has global
934 significance due to its critical role in C storage and biodiversity and relatively low levels
935 of disturbance.
- 936 ● **Impact on Indigenous and Traditional Communities:** Fires had potential to
937 significantly disrupt the lives and livelihoods of Indigenous and traditional populations,
938 exacerbating their vulnerabilities due to isolation and reduced access to resources.

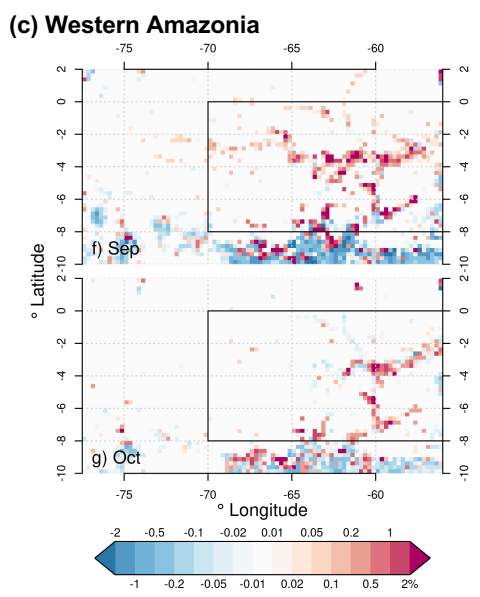
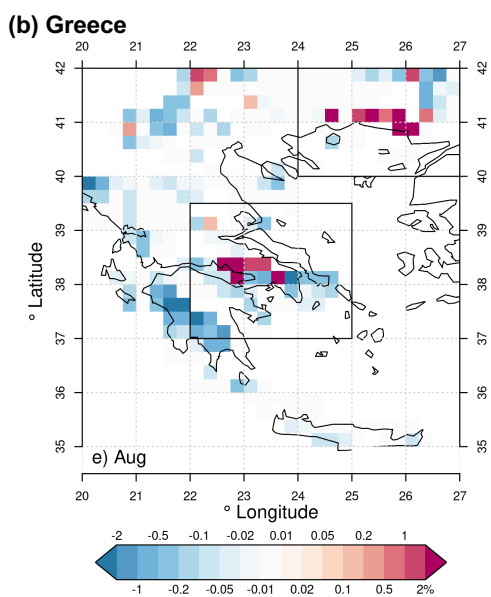
939 Abnormally high burned areas were reported in Western Amazonia during September and
940 October. **Figure 6** maps the magnitude of these anomalies. The most widespread BA
941 anomalies emerged in August 2023 and extended through to November 2023 (**Figure S11**).

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Figure 6: Spatially explicit anomalies in BA fraction (%) during key months for focal events in (a) Canada, (b) Greece and (c) Western Amazonia. Plotted data are the absolute change from the climatological mean BA fraction for the month (%), based on MODIS BA product aggregated to 0.25°. Red indicates higher BA in that month of 2023 vs the 2002-2022 climatological average for that month. Boxes indicate focus events for our analyses in this report. The top panels show anomalies in Canada for various months, the lower-left panel shows anomalies in Greece for August, and the lower-right panels show anomalies in Western Amazonia during September and August.

3 Diagnosing Drivers and Assessing Predictability

3.1 Methods

3.1.1 Predictability of Focal Extremes

We evaluated the time frame over which extreme events could have been forecasted using a common metric of fire danger, the Fire Weather Index (FWI). Developed by the Canadian Forest Service as part of the Canadian Forest Fire Danger Rating System (CFFDRS; van Wagner, 1987), the FWI comprises various components that consider the influence weather on fire danger, with 2m temperature, 10m wind speed, precipitation, and 2m relative humidity as prerequisite variables. The FWI combines three sub-indices, which are fuel moisture codes representing vegetation moisture state at different layers in the forest floor, as well as a spread index influenced by fuel moisture state and wind speed (van Wagner, 1987). A higher FWI indicates fire weather conditions more conducive to wildfires in environments with sufficient fuel load.

Owing to its original design for use in forest ecosystems, the FWI is especially useful for predicting the likelihood and severity of extreme events in ecosystems where weather is the primary limitation to fire (i.e. those mainly limited by moisture or temperature); its accuracy in forecasting BA in fuel-limited ecosystems is more limited (Carvalho et al., 2008; Bedia et al., 2015; Abatzoglou et al., 2018; Jones et al., 2022). Its applications encompass early warning systems, pre-suppression and suppression planning, prescribed burn planning and effectively alerting authorities and the public to abnormal fire danger conditions. FWI is extensively used in operational global information platforms such as the European Forest Fire Information System (EFFIS; <https://forest-fire.emergency.copernicus.eu/>, last access: 9 July 2024) and the Global Wildfire Information System (GWIS; <https://gwis.jrc.ec.europa.eu/>, last access: 9 July 2024), and the Canadian Wildland Fire Information System (CWFIS; <http://cwfis.cfs.nrcan.gc.ca/>, last access: 9 July 2024). The FWI is not the only index for fire danger, and other fire danger systems or sub-indices of this system may correlate more strongly with BA or fire behaviour metrics in some environments. Nonetheless, FWI is widely applied due to its good performance across a range of environments (Di Giuseppe, 2016; Jones et al., 2022) and so we adopt it in the current work.

In addition to well-established fire danger forecasts with lead times of a few days, skilful predictions of fire danger can be made on sub-seasonal to seasonal time scale (S2S) for Mediterranean Europe (Bedia et al., 2015), United States (Roads et al., 2010) and Asia (Spessa et al., 2015). Drought and fire weather conditions throughout the world have been found to correlate with large scale climate patterns such as the El Niño Southern Oscillation (ENSO) (Field et al., 2016; Chen et al., 2017), the Indian Ocean Dipole (Cai et al., 2009) for which current numerical weather prediction systems showcase a predictive skills. Other climate modes such as the Atlantic Multidecadal Oscillation and the Pacific Decadal Oscillation have been shown to influence fire-favourable weather conditions for some seasons and regions (Aragão et al., 2018; Turco et al., 2018) However, due to the larger uncertainties in their predictions, they are not considered here.

Following the concept of seamless prediction of fire weather on S2S timescales (Wetterhall et al., 2018; Di Giuseppe et al., 2020; Di Giuseppe et al., 2024; Dowdy, 2020), we collated FWI data from reanalyses and forecasts designed to operate on S2S lead times of 10 days to 7 months. Here, we take FWI estimates from the ERA5-Land reanalysis product (Muñoz-Sabater et al., 2021) as a proxy for observed FWI. Forecasts at different lead times are taken from the operational high-resolution ECMWF weather system, and seasonal predictions are sourced from ECMWF's seasonal forecasting system, ECMWF-SEAS5 (Johnson et al., 2019; Di Giuseppe et al., 2020; Di Giuseppe et al., 2024). A comparison between reanalysis and

1015 forecast provides an indication of how weather forecast errors translate into FWI uncertainties
1016 (predictability). Additionally, the predictions are compared to recorded peaks in fire activity,
1017 both in terms of burned areas and active fires as observed by the MODIS satellites, to provide
1018 a qualitative assessment (skill) of the correlations between landscape flammability and actual
1019 fire events.

1020
1021 The prediction systems utilised here vary in their spatial and temporal resolutions. Short to
1022 medium-range FWI forecasts (up to 10 days) are available daily at a resolution of 9 km with
1023 50 ensemble members, while the FWI seasonal forecast is available monthly at a resolution
1024 of approximately 25 km, however seasonal skill is limited to 1-2 months in normal conditions
1025 (Di Giuseppe et al., 2024). All prediction systems include a measure of uncertainties through
1026 the provision of ensemble simulation (**Figure S24, Figure S25, Figure S26**). Variance across
1027 the ensemble was previously estimated to be on the order of 10%-15% (Vitolo et al., 2020).

1028
1029 Here, the predictive skill of the models is assessed qualitatively by visually examining the
1030 extent to which extreme FWI (specifically the ensemble mean) was as a precursor of several
1031 focal events, replicating the use of this indicator by fire agencies during the fire season. The
1032 approach is designed to partially replicate the interpretation and application of the FWI by fire
1033 management agencies. Most fire agencies would have local information on fuel conditions and
1034 would thus be able to interpret FWI values in a more informed manner, reducing the
1035 dependence of decisions on FWI anomalies alone. The FWI should not be evaluated using
1036 traditional skill scores, as these would be dominated by false alarms. We maintain that the
1037 FWI is an index representing flammability and, therefore, cannot be fairly validated against fire
1038 activities.

1039

1040 **3.1.2 Identifying Key Drivers of Focal Events**

1041

1042 **3.1.2.1 Modelling Systems**

1043

1044 We used two modelling systems with similar fire predictors to diagnose the drivers of each
1045 focal event. The models are the PoF model (McNorton et al., 2024) and the ConFire attribution
1046 framework (Kelley et al., 2019; Kelley et al., 2021). The PoF model diagnoses the drivers of
1047 active fire (AF) observations from the MODIS MCD14ML active fire product (collection 6.1; 1
1048 km resolution; Giglio et al., 2016; NASA FIRMS, 2020) using Shapley values (Lundberg and
1049 Lee, 2017), while ConFire diagnoses drivers of BA from the MODIS BA product (Giglio et al.,
1050 2018; regridded to 0.5°). Fires flagged as low confidence in the AF product were not used.
1051 Although AF and BA have been used widely in global and regional scientific studies, there are
1052 substantial differences between the two branches of fire observation as reviewed extensively
1053 elsewhere (Roy et al., 2008; Di Giuseppe et al., 2021; Chuvieco et al., 2019), and the strength
1054 of the relationship between them can vary regionally (van der Werf et al., 2017; Hantson et
1055 al., 2013). Our use of two observational fire products and two distinct model approaches
1056 provides a way to account for inherent uncertainties in the observability of different fire events
1057 and the uncertainties in the methodologies.

1058

1059 Both modelling approaches use a number of individual predictors of AF or BA, which we refer
1060 to as ‘drivers’. The drivers are grouped into four main categories, which we refer to as the
1061 controls. PoF and ConFire both include weather, fuel abundance, and fuel moisture as controls
1062 on fire. In addition, PoF (but not ConFire) includes an “Other” category of controls and ConFire
1063 (but not PoF) includes a “human” category of controls, as per **Table 3**. PoF drivers in “Other”
1064 include ignition and suppression effects as well as the residual error between predicted and
1065 observed fire activity. Grouping the set of drivers between the four identified controls —
1066 weather, fuel moisture, fuel load, and human/other— is not always straightforward, as fuel
1067 moisture and weather variables are strongly correlated, and fuel load is also related to weather
1068 conditions. Hence, some drivers can be associated with more than one control (**Table 3**). The

1069 categorisation stems mostly from the way the driving datasets have been obtained and their
1070 underlying resolutions. We have also considered the traditional approach of assessing fire
1071 weather in isolation within most fire danger assessment metrics. We believe that grouping
1072 these metrics under the umbrella term 'control weather' offers a concise way to reference the
1073 drivers of the fire weather index (Matthews, 2009). Despite this, it is important to note that the
1074 techniques employed ensure contribution from specific variables cannot be double-counted
1075 between categories. Both ConFire and PoF are capable of disentangling the contributions of
1076 individual drivers within the same control category (for example, the separate contributions of
1077 dead or live vegetation) and quantifying these contributions (Kelley et al., 2019; McNorton et
1078 al., 2024). However, we will focus our analysis on the impact of the controls.

1079
1080

1081 **3.1.2.2 Drivers and Controls Used in Fire Event Analysis**

1082

1083 For our assessment of the contribution of weather and fuel moisture to the anomalous events
1084 we take several predictors from ERA5-Land (9 km resolution; Muñoz-Sabater et al., 2021),
1085 specifically variables that are known to correlate with AF or BA (Bistinas et al., 2014; Haas et
1086 al., 2022). The drivers considered for each control are listed in **Table 3**. For the weather
1087 component in isolation, we use 2m temperature, 2m dewpoint temperature, 10m wind speed
1088 and daily total precipitation (note that these are the prerequisite variables used in the
1089 formulation of the FWI; van Wagner, 1987). We use a fuel characteristic model to estimate the
1090 fuel load and fuel moisture components following McNorton and Di Giuseppe (2024), with
1091 model estimates of fuel moisture constrained by estimates of leaf area index (LAI) from the
1092 ECMWF's Integrated Forecast System (IFS) and model estimates of fuel loads constrained
1093 by aboveground biomass estimates from the ESA CCI (Santoro and Cartus, 2021) and net
1094 ecosystem exchange estimates from the Copernicus Atmosphere Monitoring Service (CAMS;
1095 Agustí-Panareda et al., 2019). Additional predictors regarding fuel load and state include
1096 vegetation cover and type (**Table 3**). Proxies for ignition and suppression controls, placed
1097 within the "Other" set of controls, are more challenging to establish. Currently, we use
1098 population density, urban fraction, cropland fraction, pasture fraction, lightning, orography
1099 (**Table 3**). For consistency all variables are interpolated to 9 km resolution. 'Others' not only
1100 include factors related to ignitions but also the fraction of predictions missed by the models.
1101 This is important because this category weights the importance of unaccounted-for factors.

1102

1103 Another important aspect is that models do not assume a specific direction for each factor's
1104 influence on fire activity. Consequently, wetness can correlate with increased fire likelihood in
1105 some locations and reduced fire likelihood in others. This aligns with established theory in our
1106 field: in fuel-limited regions where grass and herbaceous fuels dominate, high rainfall
1107 promotes fuel accumulation and increases fire extent. Conversely, in fuel-rich regions with
1108 high tree cover, high rainfall increases fuel moisture and reduces fire extent. See "**Modelling
1109 frameworks**" in the **Extended Methods Supplement** for a detailed description. See
1110 "**Evaluation**" in the **Extended Methods Supplement** for detailed evaluation.

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Table 3: Drivers of fire and their parent control group included in the event fire analyses using ConFire and PoF. Drivers are individual variables, which serve as proxies for the influence of weather, fuel load, fuel moisture, or other controls on BA. ** The ‘Other’ category includes proxies for ignition and suppression controls plus the missed prediction. Note that for ConFire, explanatory variables can be associated with multiple controls (Kelley et al. 2019). Positive (+ive) or Negative (-ive) under “ConFire control” describes if a driver increases or decreases BA in ConFire.

Driver	POF control	ConFire controls	Frequen cy	Time Period	Source	Reference
2m Temperature	Weather	Weather +ive	Daily	Jan 2014-NRT	ERA5-Land	Muñoz-Sabater et al. 2021
2m Dewpoint Temperature	Weather	Weather -ive	Daily	Jan 2014-NRT	ERA5-Land	Muñoz-Sabater et al. 2021
10m Wind Speed+	Weather	Not used	Daily	Jan 2014-NRT	ERA5-Land	Muñoz-Sabater et al. 2021
Precipitation	Weather	Weather -ive	Daily	Jan 2014-NRT	ERA5-Land	Muñoz-Sabater et al. 2021
Live Leaf Fuel Load	Fuel Load	Not used	Daily	2014- 2020	Fuel Model	McNorton et al. 2024a
Live Wood Fuel Load	Fuel Load	Not used	Daily	2014- 2020	Fuel Model	McNorton et al. 2024a
Dead Foliage Fuel Load	Fuel Load	Not used	Daily	2014- 2020	Fuel Model	McNorton et al. 2024a
Dead Wood Fuel Load	Fuel Load	Not used	Daily	2014- 2020	Fuel Model	McNorton et al. 2024a
Mean & Max Vegetation Optical Depth (VOD) of the last 12 months	Not used	Fuel Load +ive	Monthly	2014-NRT	Satellite-Soil Moisture and Ocean Salinity (SMOS)	Wigneron et al 2021
Vegetation Optical Depth (VOD)	Moisture	Moisture -ive	Monthly	2014-NRT	Satellite (SMOS)	Wigneron et al 2021
Low Vegetation (LAI)	Fuel Load/Moisture	Not used	Monthly	2002-2020 climatology	Satellite (multi-sensor)	Boussetta et al., 2021
High Vegetation (LAI)	Fuel Load/Moisture	Not used	Monthly	2002-2020 climatology	Satellite (multi-sensor)	Boussetta et al., 2021
Live Fuel Moisture Content	Fuel Moisture	Fuel Moisture -ive	Daily	2014-NRT	Fuel Model	McNorton et al. 2024a
Dead Fuel Moisture Content	Fuel Moisture	Fuel Moisture -ive	Daily	2014-NRT	Fuel Model	McNorton et al. 2024a
Snow cover	Fuel Moisture	Snow -ive	Daily	2014-NRT	ERA5-Land	Muñoz-Sabater et al. 2021
Pasture	Not used	Ignitions +ive Suppression -ive	Annual	2014-2023	HYDE	Klein Goldewijk et al., 2011
Cropland	Not used	Ignitions +ive Suppression -ive	Annual	2014-2023	HYDE	Klein Goldewijk et al., 2011
Urban population	Not used	Ignitions +ive Suppression -ive	Annual	2014-2023	HYDE	Klein Goldewijk et al., 2011
Rural populations	Not used	Ignitions +ive Suppression -ive	Annual	2014-2023	HYDE	Klein Goldewijk et al., 2011
Lightning	Not used	Ignitions +ive	Monthly	2000- 2020 climatology	LIS/OTD	Cecil et al., 2014
Type of Vegetation	Other**	Not used	Fixed	Jan 2014-NRT	ECLand	Boussetta et al., 2021
Urban Fraction	Other**	Not used	Fixed	Jan 2014-NRT	ECLand	McNorton et al. 2023
Orography	Other**	Not used	Fixed	Jan 2014-NRT	ECLand	Boussetta et al., 2021

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3.1.2.3 Analysis of Fire Drivers

The PoF system uses gradient-boosted decision trees from the XGBoost library on detected AF (McNorton et al., 2024). The training iteratively adds decision trees to an ensemble of models to correct for errors made by previous iterations, resulting in a computationally efficient optimization (Chen and Guestrin, 2016). The system training uses a classifier approach which defines a positive hit as an AF detection within the grid cell on a given day. The driver attribution is performed using the SHapley Additive exPlanations (SHAP) method taken from the SHAP library (Lundberg and Lee, 2017). These are then combined to provide overall attribution to one of the four controls for AF predictions.

ConFire uses Bayesian Inference to assess fire behaviour uncertainty by evaluating control strengths' impact on BA. Instead of a single output, it produces a probability distribution based on data into a simplified fire model. Variables like weather and fuel moisture contribute simultaneously to the prediction of monthly BA. Monthly averages of daily values are aggregated using the Fogo Local Analisado pela Máxima Entropia (FLAME) system (Barbosa, 2024). Each control combines drivers using logistic functions. Bayesian inference optimises driver contributions and accounts for stochasticity, capturing fire unpredictability under similar conditions (Kelley et al., 2021). The mean logit-transformed BA distance with and without this term measures additional uncertainty. The model is trained and ran from 2014 to 2023 using driving datasets common across this period. Initially trained on 50% of BA, it is then applied predictively to the full dataset. A separate evaluation is conducted, training on data from 2014 to 2018 and testing against 2019 to 2023, following the protocol employed by Barbosa (2024). See **“Modelling frameworks”** and **“Evaluation”** in the **Extended Methods Supplement** for a detailed description.

For both models we include an estimate of uncertainty. ConFire is designed as an uncertainty quantification model, providing BA probabilities and their likelihoods for each region. Results in this report are based on 5-95% confidence intervals or likelihoods, supporting central (median) estimates. ConFire uses Bayesian inference to quantify how various factors impact fire occurrence, assuming that fuel increases BA, moisture decreases it, ignitions increase it, and suppression decreases it. The model trains on historical data to determine influence levels and accounts for uncertainties from fire stochasticity and unconsidered factors. Its probability distribution is a logit zero-inflated function, assessing changes in extreme fires even with small observed areas. ConFire manages uncertainties from unpredictable weather and vegetation responses. ConFire quantifies uncertainty estimates from different drivers by constraining them with observations and addresses structural uncertainties, such as missing explanatory variables and errors in mechanistic relationships. While it represents one relationship of control to BA, the probability distribution accounts for uncertainty across various potential relationships. Noise is considered, with the stochasticity of burned area accounted for both in areas with no fire and in the probability of different potential levels of burning where fire does occur. We test ConFire's uncertainty quantification using Bayesian inference evaluation techniques. However, ConFire does not account for some uncertainties, such as potential changes in feedbacks between fire and vegetation when transitioning to a new fire regime out of sample. The model assumes the accuracy of one bias-corrected model for fuel load (JULES-ES), neglecting uncertainty from Dynamic Global Vegetation Models (DGVM). While BA anomalies are used to reduce observational bias, any disagreements in bias across space or time are not included in ConFire's training.

Meanwhile, PoF outputs are provided in terms of probabilities calculated using ensemble predictions from weather forecasts each to generate a set of binary classifiers. The probabilities are therefore based on a wide parameter space, taking into account uncertainties in both the input parameters and the stochasticity of the classification algorithm itself.

1176 **3.2 Results**

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1178 **3.2.1 Predictability of Focal Events using Fire Weather Forecasts**

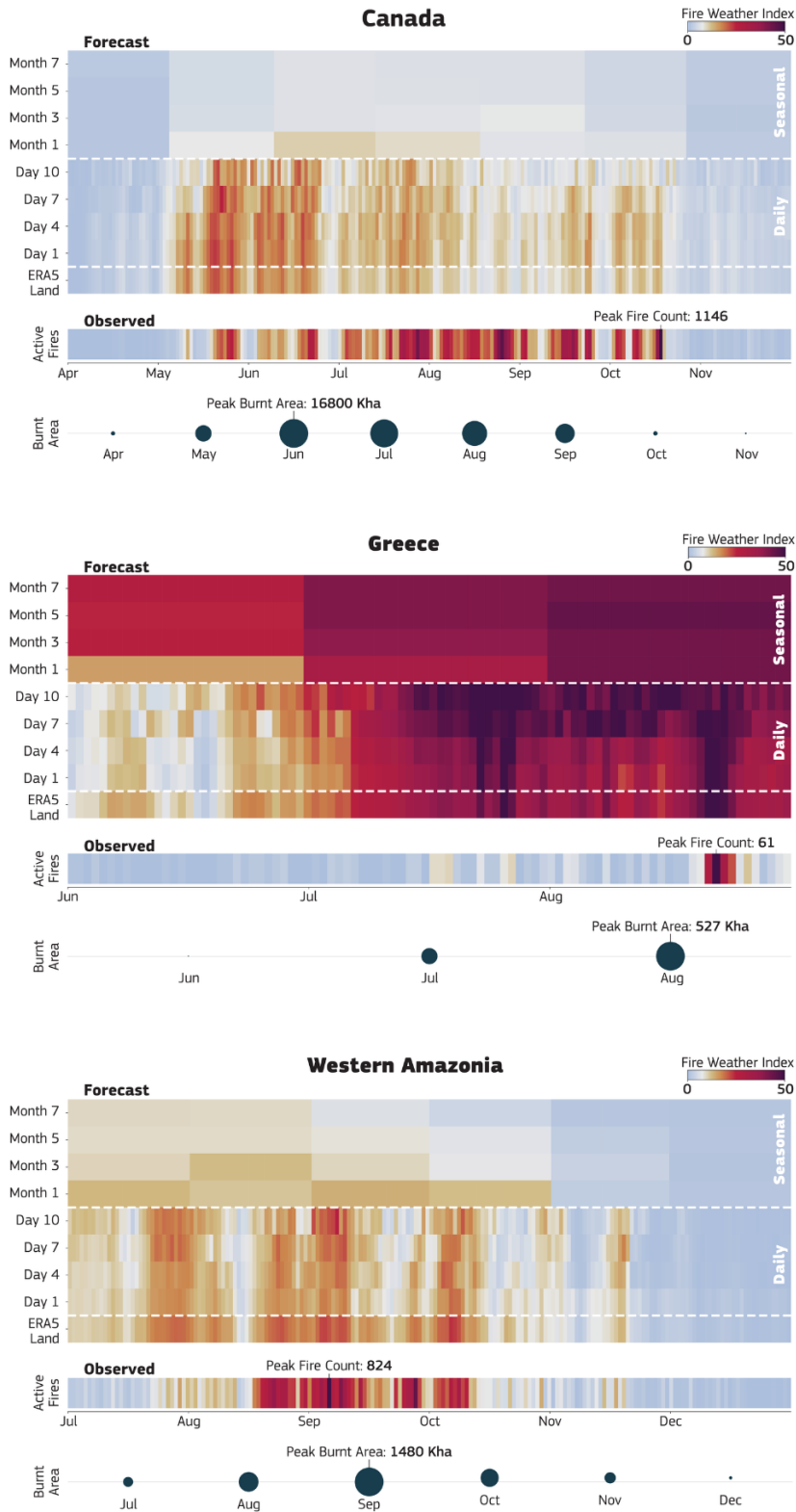
1179

1180 **3.2.1.1 Canada**

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1182 The early establishment of fire weather conditions as well as the late cessation of the fire
1183 season are well captured in the FWI reanalysis in Canada (**Figure 7**). The FWI also captures
1184 the intermittent pattern of fire danger and its correlation to actual fire activity. However, at the
1185 seasonal time scale, the signal is weakened, and there were no prior indications that the
1186 upcoming season would have been as extreme as it was with respect to fire activity (**Figure**
1187 **7**). For most part of 2023 Canada was in drought. The weather-limited nature of the Canadian
1188 fire season means that the FWI modelling framework serves as an essential indicator of
1189 anomalous conditions, acting as a prerequisite for the intensity and spread of fires. It provides
1190 valuable insights into the sequence and extent of extreme fire weather days during such
1191 events. Notably, in this region, peaks of fire activity correspond to peaks of landscape
1192 flammability, and there is a good correlation between observed fire activity and predicted fire
1193 danger.

Fire Activity & Prediction 2023



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Figure 7: Chicklet plots displaying seamless FWI predictions over time from various forecasting systems of the ECMWF (see Methods). The x-axis corresponds to specific dates throughout the year, while the y-axis denotes either observations or the time leading up to the

1198 date when a forecast was generated. The vertical colour coherence allows for quick
1199 identification of the time windows of predictability associated to the observed fire activity both
1200 provided in terms of number of detected active in a day fires and total burned areas in a month
1201 (circles).

1202

1203 **3.2.1.2 Greece**

1204

1205 The establishment of fire-prone conditions in the Mediterranean, particularly in Greece, is part
1206 of the region's seasonal weather cycle (**Figure 7**). In 2023, this pattern persisted, and extreme
1207 landscape flammability could be forecasted well in advance. In arid and semi-arid regions fire
1208 occurrence is driven not solely by weather but also by fuel availability and its intermittent short-
1209 term drying. In these regions the FWI often reaches extreme levels for much of the summer.
1210 However, fires do not always occur even when the FWI is extreme, as ignition and early
1211 suppression play a crucial role. The anomalous fire extent in Alexandroupolis, including the
1212 large Evros fire, highlights the limitations of relying solely on fire weather indices in these
1213 areas. There were no discernible indications in the FWI records that the particular day was
1214 more extreme than the days before or afterwards, emphasising the need for a more holistic
1215 approach to fire risk assessment in regions where fuel is a limiting factor or live fuel moisture
1216 plays a crucial role in the extent of the fires (Di Giuseppe, 2023).

1217

1218 **3.2.1.3 Western Amazonia**

1219

1220 The correlation between FWI and fire activity in the western Amazonia region at the shorter
1221 lead times is generally poor, primarily due to the strong dependency on either lightning or
1222 human ignitions (Kelley et al., 2021). In 2023, this pattern persisted (**Figure 7**), with the onset
1223 of fire weather following the seasonal pattern well ahead of the time where fires were triggered.
1224 Seasonal predictions indicated high fire danger during the summer period probably driven by
1225 El Niño conditions.

1226

1227 **3.2.2 Identifying Key Drivers of Focal Events**

1228

1229 Weather, fuel moisture, fuel abundance, and ignitions are the four primary controls identified
1230 as influencing the occurrence and intensity of the focal fire events. Anomalies in individual
1231 drivers of these controls, such as temperature or soil moisture, are calculated by comparing
1232 regional daily 2023 averages with the average for 2003-2022. Dead fuel, with its lower
1233 moisture content and higher combustibility, often plays a significant role in determining fire
1234 ignition. During extreme events, it is the dry live fuel that burns, contributing to the overall
1235 severity and intensity of the fire.

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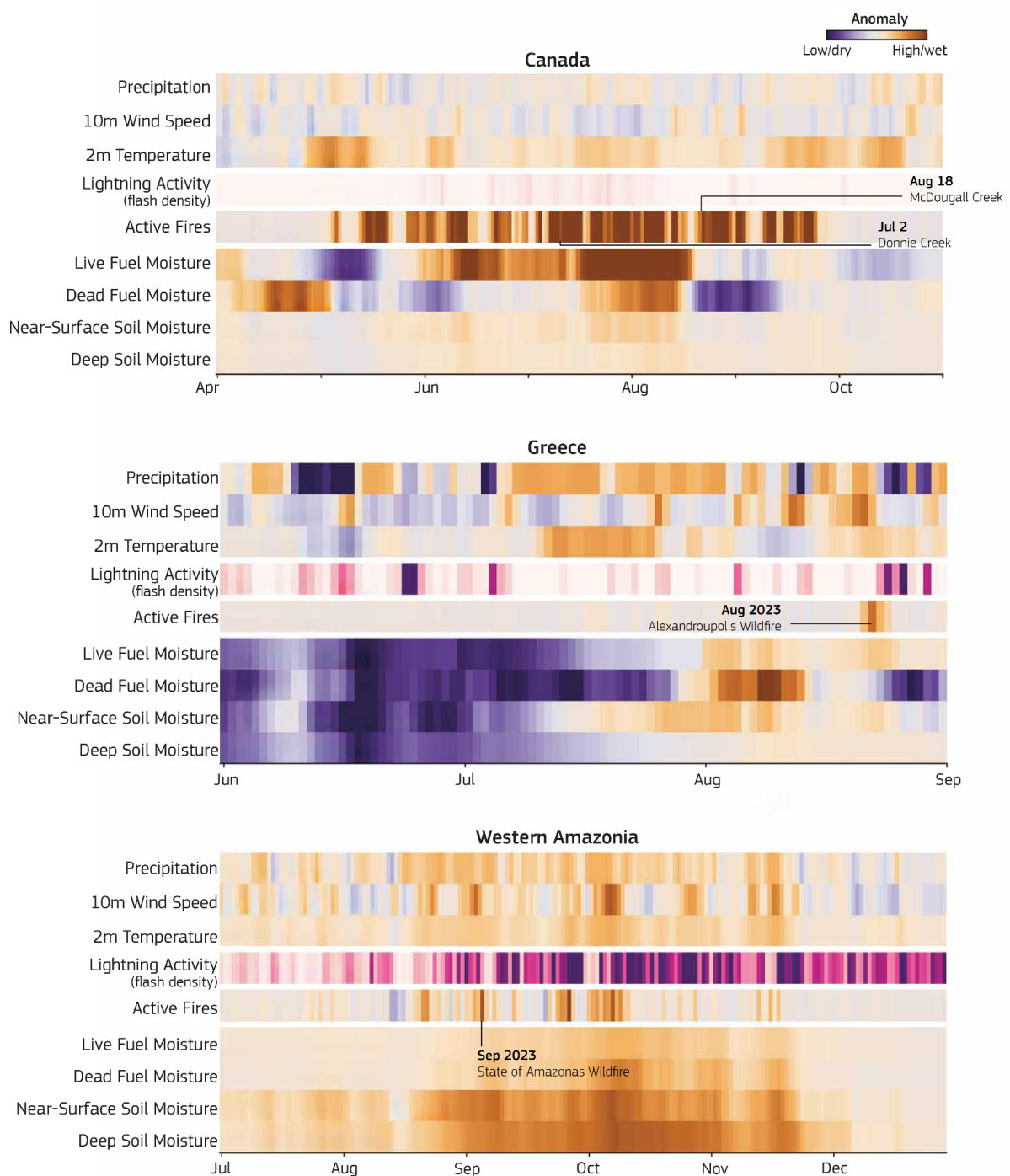
1237 Analysing the time series of key drivers contextualises the conditions under which events
1238 occurred (see **Figure 8**). However, leveraging the PoF and ConFire models allows for a
1239 statistical causality attribution of the four controls for observed fire occurrence (see **Figure 9**).
1240 These models will provide control attribution even if no fire event is recorded, with a low
1241 probability across all controls indicating an accurate prediction. High fire probability without
1242 recorded fire activity could indicate successful suppression or fire-prevention policies.
1243 Unaccounted human influence is categorised under "other," encompassing variables not
1244 forecasted by the models. This analysis enhances our understanding of fire activity controls
1245 and helps identify missing information that degrades the quality of the prediction.

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Fire drivers in 2023



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Figure 8: Anomaly driver stripes for the three focal events. The drivers are selected to contextualise the conditions under which the examined events took place. All values are expressed as anomalies compared to the 2003-2021 climatology with the exception of lightning activity which is expressed as absolute flash density.

1257 **3.2.3 Drivers of Active Fire Extremes**

1258

1259 **3.2.3.1 Canada**

1260

1261 Persistent fire-favourable weather conditions played a crucial role in controlling the extent of
1262 active fires in Canada during the summer of 2023. Dry weather contributed to extensive drying
1263 of both live and dead vegetation, further exacerbating fire risk (**Figure 8, 13**). Most of the
1264 explainability of the Canada event comes from anomalous weather conditions. Increased
1265 lightning activity often coincides with or precedes significant fire periods, indicating lightning
1266 as a key source of ignitions in the region given the contribution of the cloud-to-ground flashes
1267 to the total predicted lightning activity. This is in agreement with the attribution of 59% of
1268 wildfires and 93% of total BA to lightning ignition sources in Canada during 2023 (Jain et al.,
1269 2024). Adverse weather conditions in mid-May in western Canada were identified as influential
1270 factors in shaping fire events. However, multiple instances of intense burning events, notably
1271 in mid-July, early August, and late September, fall into the 'Other' category, heavily
1272 contributing to the total number of events for which there is no attribution among the controls.
1273 The fact that clusters of events were not predicted suggests potential inadequacies in
1274 accounting for some ignition sources or accurately representing fire propagation across these
1275 landscapes.

1276

1277 **3.2.3.2 Greece**

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1279 The driver anomalies (**Figure 8**) and control attribution (**Figure 9**) did not suggest an
1280 abnormally fire-prone year in Greece, failing to explain the large fire extent around the time of
1281 the Evros fire near Alexandroupolis. An anomalously wet spring may have led to increased
1282 foliage and subsequently quick drying of plant material. A sustained dry period in late July and
1283 August further dried out new foliage, creating favourable conditions for fire activity, as
1284 indicated by the anomalously dry live and dead fuel moisture content in August. Despite these
1285 conditions, the unexpected extent and severity of fires around Alexandroupolis were not
1286 predicted, highlighting the intrinsic difficulties in forecasting isolated extreme events even
1287 when most predictors are included. Additionally, the high wind speeds at the time partially
1288 contributed to the extensive BA during the fire.

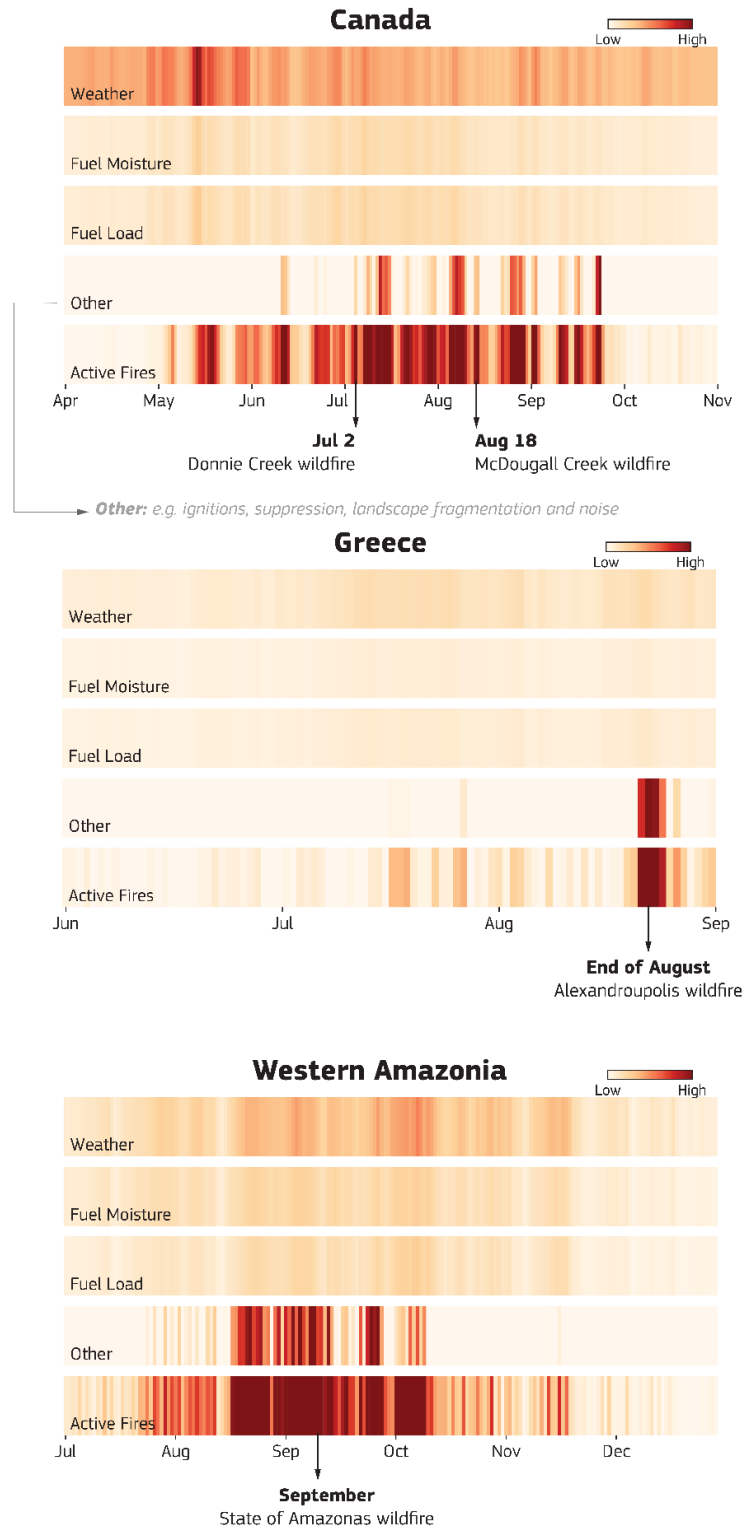
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1290 **3.2.3.3 Western Amazonia**

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1292 Prolonged drought conditions driven by a positive ENSO (Aragão et al., 2007; Jiménez-Muñoz
1293 et al., 2016), stemming from anomalously low rainfall and high temperatures, created
1294 favourable conditions for an active fire season in Western Amazonia (**Figure 8, Figure 9**).
1295 These conditions had a significant impact on the typically wet ecosystem, affecting soil
1296 moisture as well as live and dead fuel moisture. Despite weather conditions serving as a
1297 persistent control for fire activity, several intense active fire periods in late August and
1298 throughout September were not predicted, possibly due to unrepresented ignition sources.
1299 Additionally, fire activity from September onwards was intensified by intense lightning activity,
1300 characteristic of the region, which substantially contributed to ignitions.

Fire controls 2023



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Figure 9: Contributions of different fire controls to daily active fire anomalies in our focal events. All values are scaled to the observed daily fire anomaly such that the sum of the 4 daily control values matches the total observed anomaly.

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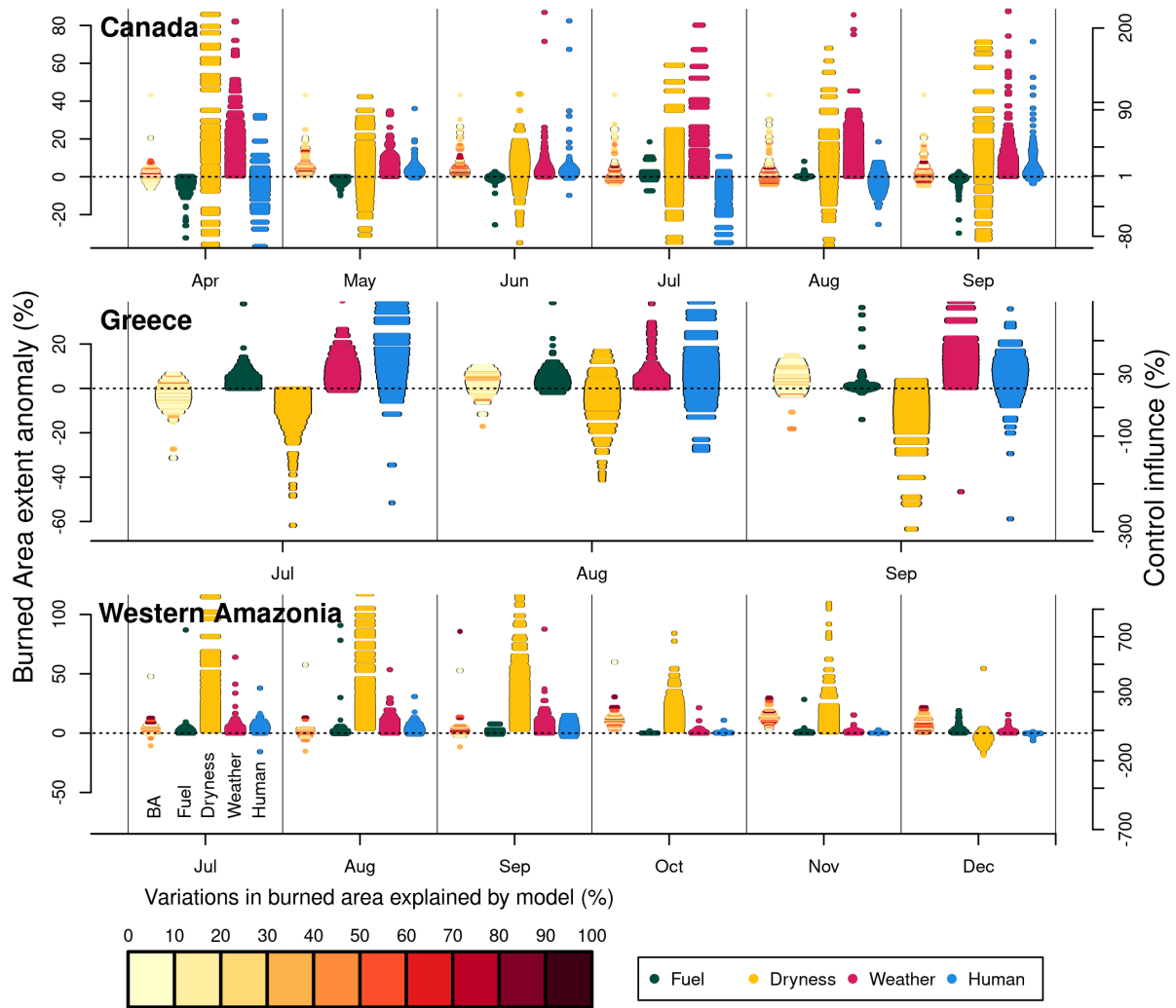
3.2.4 Drivers of Regional Burned Area Extremes

3.2.4.1 Canada

ConFire detected a significant anomaly in BA starting in late April, as seen in the observations (**Figure S9**) with very high confidence between May (99.2% likelihood) and June 2023 (>99.9% likelihood; **Figure 10**). While less confident in a positive anomaly in September and August (71% likelihood), ConFire detects the possibility of much higher burning in August and September, corresponding with the increase in burning in the Western Shield (**Figure S9**). **Figure 10** shows the controls that contribute to these anomalies. Our analysis indicates a >99.9% likelihood that elevated fire weather conditions persisting throughout the 2023 fire season, led to a notable increase in burning, explaining 19 [4.6-45]% of the BA anomaly in May and 13 [1-110]% in June (median estimates, with 5th-95th percentile range in square brackets). Anomalous weather conditions subsided in May through the early summer, though by September showed an increased likelihood of contributing to the increase in BA anomaly seen in the late fire season. Drier fuel conditions could have contributed significantly to the increase in BA (up to 65% of the BA anomaly in May and 45% in June), though with low confidence (60.5% in May and 61.3% in June), and wetter fuels exerting a suppressive effect on fire spread was also possible, suggesting their potential role in mitigating fire severity. A small but confident suppressive effect (100% likelihood in May, 64.8% likelihood in June) from fuel load was observed, reducing relative increases in BA by 1.4 [0.17-7.1]% in May. Direct human-induced landscape changes exhibited a small impact on the extent of burned areas (likelihood of 97.4% in May), explaining between 5.4 [1.2-22]% of the anomaly in BA in May and 5.2 [0.6-24]% in June.

3.2.4.2 Greece

The analysis reveals an anomalously high BA, particularly from mid-August onwards, though with a lower confidence level compared to the Canadian case (69.9% likelihood; **Figure 10**). **Figure 10** shows the controls that contribute to these anomalies. Our findings demonstrate with very high confidence the presence of anomalously high fire weather conditions during the 2023 fire season in Greece (98.3% in July, and >99.9% in August and September). In August, these conditions explained 24 [4.3-140]% of the increased BA, increasing to 34 [5.6-170]% by September. Assessing the impact of fuel moisture on BA, our analysis shows a wide range of possibilities, with confidence ranging from a 21% increase to 180% decrease in relative BA extent, which would have offset some of the increases from fire weather. This uncertainty underscores the complexity of the interactions between fuel moisture and fire behaviour in Greece. Although direct human-induced landscape changes exerted greater influence on BA extent in Greece than in the other focal regions, this influence remained small compared with weather factors. While the analysis indicates a slightly higher than normal fuel load with only a low likelihood of having a substantial influence on increased levels of burning.



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Figure 10: Influence of fuel load, fuel dryness, fire weather, and human controls on burned area (BA) anomalies in 2023. **(Top row)** The colour indicates the modelled BA anomalies as relative increases/decreases in BA extent (% relative to the climatology). **(Rows 2-4)** The % of the anomaly explained by each control. At each timestep, the gradient of colours along the y-axis represent the range of modelled outcomes: high diversity in the colours indicates high uncertainty, whereas a consistent colour (e.g., red for positive BA anomalies) indicates greater confidence in the projected direction and magnitude of change. Dotted black line indicated threshold between positive and negative anomaly. **(Bottom row)** Uncertainty measure representing potential variations around our simulated BA anomaly as a result of factors not considered by the modelling framework, as a fraction of land area. Lower numbers (lighter colours) indicate higher confidence in the BA anomaly in the top row.

1362 **3.2.4.3 Western Amazonia**

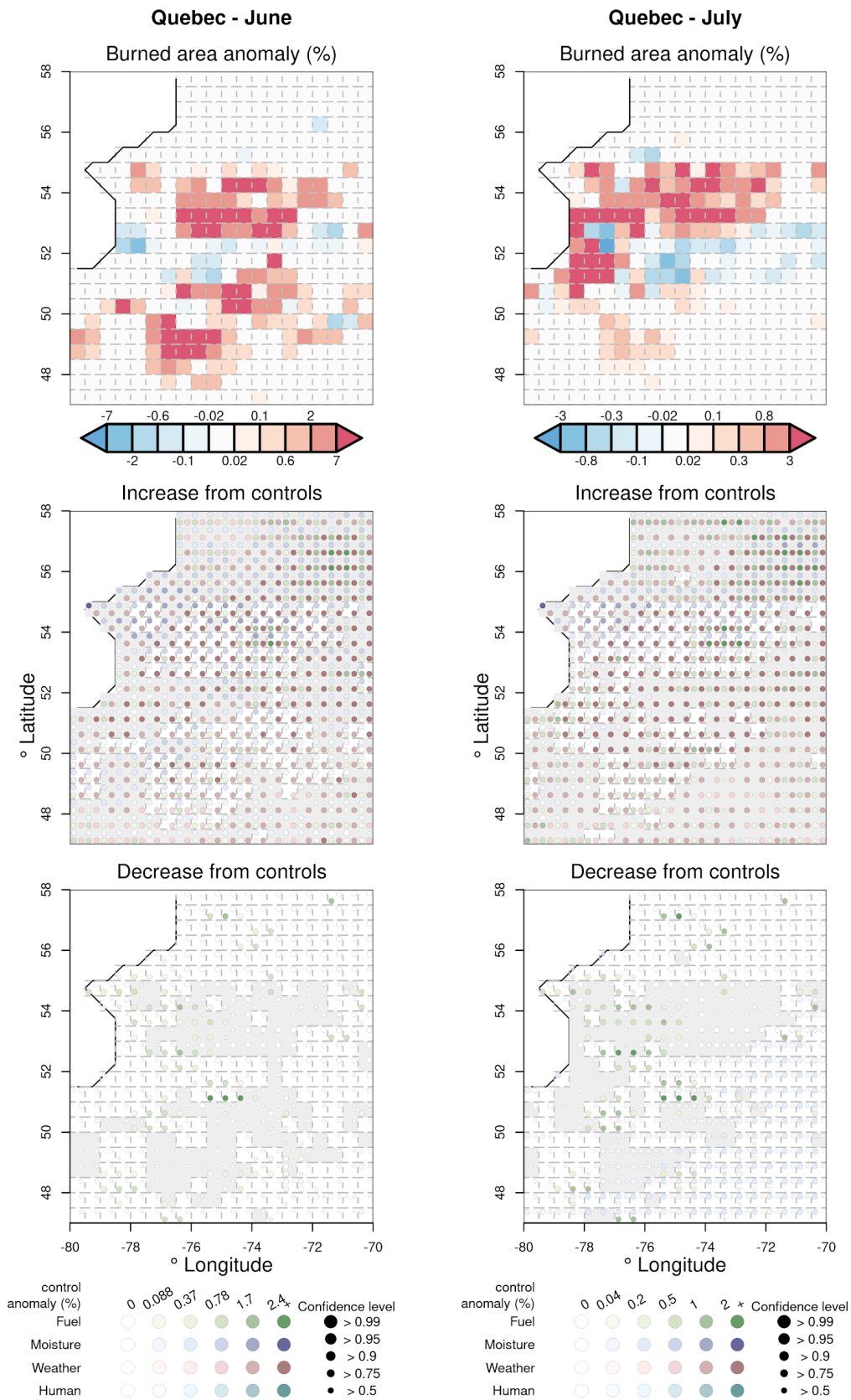
1363
1364 Our analysis indicates a reasonably high confidence that the considered drivers contributed
1365 to anomalous high burning during September (73.8% likelihood), October (94.9% likelihood)
1366 and November (96.1% likelihood; **Figure 10**). **Figure 10** shows the controls that contribute to
1367 these anomalies. The primary driver of the observed BA anomaly appears to be dry fuel
1368 conditions, with a very high likelihood (99.7%) of drier-than-normal conditions persisting
1369 through November. This led to a substantial increase in BA, explaining at least 57% of the
1370 increase in BA. While fire weather conditions were also elevated (likelihood >99.9%), their
1371 impact on BA was comparatively lower, resulting in at least 2% increase in BA during October
1372 and November, though with a small probability of contributing much more. Direct human
1373 influence was identified as a contributing factor, with a high likelihood (92.9% in September,
1374 94.5% in October) of increasing burned area. However, the magnitude of this influence was
1375 approximately one-tenth of that attributed to fuel dryness. There is little confidence in the
1376 direction of the effect on BA anomaly, with potential influences ranging from a slight
1377 suppressive effect (26.5% likelihood) to potentially explaining the majority of the increased BA
1378 in September, to virtually no impact in October. This suggests that fuel dynamics played a
1379 minor role in driving the observed fire activity. The analysis reveals a higher confidence in the
1380 simulations indicating positive anomalies, indicating a robust signal in the attribution of drivers
1381 to observed BA anomalies.
1382

1383 **3.2.5 Spatial Variation in Drivers of Burned Area Extremes**

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1385 **3.2.5.1 Canada**

1386
1387 In Canada, most BA anomalies were linked to widespread high fire weather, with 95% of the
1388 country being influenced by higher-than-normal fire weather (**Figure S12**). There was a
1389 tendency for fuels to dry out, although this was not as widespread. Fuel load anomalies were
1390 more scattered, but areas of low fuel anomaly did correspond to some boundaries in fire
1391 extremes (**Figure 11, S12**). Increased human influence may have had some influence at
1392 suppressing fires, but this is not significant, and in some places, the model indicates a small
1393 possibility of increased fire from human activity. In the Eastern Shield, fire extremes in some
1394 areas were driven by high fire weather and dry fuel, compounded with more vegetation cover
1395 and hence higher-than-normal fuel load in some places (**Figure 11**). However, the borders of
1396 extreme fires corresponded to a suppressive effect from decreased fuel load. In the Western
1397 Shield, dry fuel and high fire weather drove fire incidents, with high fire weather dominating in
1398 some areas. Increased suppression may have had some influence at suppressing fires, but
1399 this is not significant, and in some places, the model indicates a small possibility of increased
1400 fire from human activity.
1401

1402 In June, there were high anomalous burned areas in Quebec's Eastern Shield, which were
1403 divided into two major fire components with a slightly reduced BA in between (**Figure 11**).
1404 Both components were associated with high fire weather, but in areas where high fire weather
1405 occurred without the contribution of other controls, it tended not to cause high levels of burning.
1406 The highest burned areas were mainly found in the northern component and were associated
1407 with anomalously low levels of moisture and high fire weather. Some cells with the very highest
1408 BA also showed anomalously high fuel load (**Figure 11**). In a region further north, around 56-
1409 57 degrees north and 72-80 degrees west, there was high fuel load, dry conditions, and high
1410 fire weather, but fire in the area was found to be highly fuel limited and largely insensitive to
1411 even large changes in controls (Kelley et al. 2019). The southern component of high burning
1412 corresponded with high fire weather and either fire fuel or high fuel load. Additionally, any
1413 boundaries between higher and normal/lower levels of BA also saw lower than average fuel
1414 loads, which may have inhibited fire spread (**Figure 11**).
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Figure 11: Anomalies in controls during months and regions of high BA in Quebec. The top map of each region shows BA anomalies at 0.5 degrees for that month in 2023 versus the 2014-2023 monthly average. The middle maps looks at anomalies in controls that would cause

1421 higher BA, with areas not greyed out representing regions with greater than monthly average
1422 BA in 2023. The bottom map shows drivers that would have led to lower than normal levels of
1423 burning, with areas not greyed out showing lower or non-change from monthly average BA in
1424 2023. Each grid cell has four points: green points show anomalies in fuel load, purple in fuel
1425 moisture, red in fire weather, and cyan in humans. This way, we can see if controls acted in
1426 unison to cause extreme levels of burning or prevent extreme fires from extending further. The
1427 shade of the point shows the most likely expected level of anomaly in that control, while the
1428 size shows how confident we are in the direction of the anomaly.
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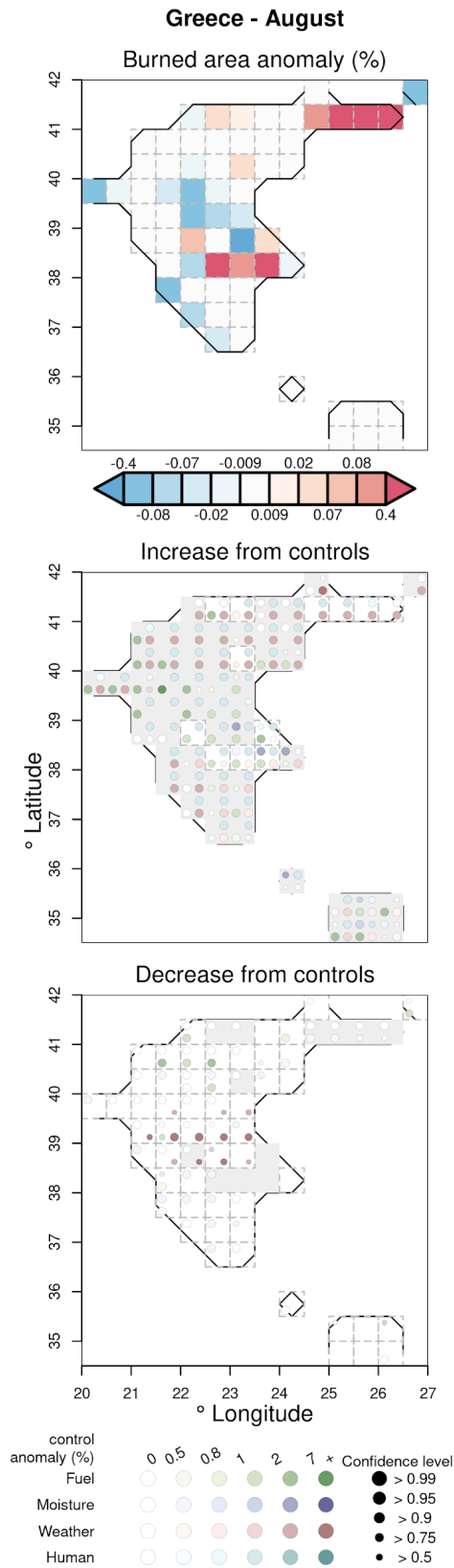
1431 In May, the Western Shield and Taiga/Boreal plains experienced higher-than-normal fire
1432 weather across the region (**Figure S13**). The increased burned areas in the west were due to
1433 extreme low fuel moisture and high fire weather, as well as higher fuel loads. In contrast, areas
1434 to the south experienced high fire weather without the extreme burned areas. Additionally,
1435 anomalies in fuel loads and burning levels became more evident in September, with some
1436 areas displaying lower-than-average fuel and burning (**Figure S13**). These anomalies
1437 persisted, with regions still experiencing high fire weather and variations in fuel moisture
1438 levels. Furthermore, the eastern areas with higher fire weather also showed higher fuel loads,
1439 while drier fuel moisture was observed in less extreme regions to the east. Additionally,
1440 specific locations saw higher fire weather and above-average fuel moisture, while areas just
1441 north of the extreme fires experienced wetter-than-normal fuel moisture (**Figure S13**).
1442

1443 3.2.5.2 Greece

1444

1445 Interestingly, most of Greece showed a tendency towards less suppression from people
1446 (**Figure S12**). However, the dominant driver over most (73%) of Greece was high fire weather,
1447 with some areas in central Greece showing notably low fire weather. These areas do not
1448 correspond to a fire anomaly, though higher fuel loads were detected. Except for central
1449 Greece, other areas of lower than average BA correspond to lower than average fuel loads
1450 (**Figure 12**). There was no significant anomaly in fuel moisture across Greece, though the
1451 northeastern fire extreme does correspond to a joint increase in fire weather and decrease in
1452 fuel moisture. Extremes in Eastern Coastal Greece correspond to anomalies in fuel load, fuel
1453 dryness, and heightened fire weather (**Figure 12**).
1454

1455 In August, Northern Greece experienced high fire weather and low fuel moisture, particularly
1456 around Alexandroupolis in Macedonia and Thrace (**Figure 12**). The region extended further
1457 east, experiencing extreme fires that reached into Central Macedonia. Unlike Canada, the
1458 framework in North Greece did not show the same level of detail in the boundaries around
1459 extreme levels of burning. However, the transition to less burning in the South of West
1460 Macedonia did correspond to reduced fuel load. Areas around Athens and Central Greece
1461 that saw unusually high levels of burning also experienced decreased fuel moisture and
1462 increased fuel load (**Figure 12**). In contrast, areas in Southern Thessaly and Central Greece
1463 that did not experience high burned areas saw lower than normal fire weather. In the
1464 Peloponnese region, there was either high fire weather or reduced fuel moisture, but these
1465 conditions rarely occurred together, which might explain the lack of increased BA throughout
1466 the region. (**Figure 12**)



1467

1468 **Figure 12:** Same as Figure 11 but for Greece.

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1470 3.2.5.3 Western Amazonia

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1472 High fire weather anomalies were almost universal across our Western Amazonia region
1473 (**Figure S12**). However, anomalies in other controls varied widely across the region, which
1474 appears to modulate the occurrence of areas with above average BA (**Figure S12, S14**). In
1475 September, the regions of high burning in the Western Amazon all exhibited higher than
1476 average fire weather, with the highest BA anomalies associated with the highest fire weathers.
1477 The areas with the highest BA anomalies along the Amazon River were located around
1478 Manaus and also showed lower than average fuel moisture and higher than average fuel load.
1479 The very highest pixels exhibited the anomaly in increased fuel loads. In the southern and
1480 central part of the region (65 to 62 west/7/8 south), near Porto Velho, BA anomalies were
1481 associated with extremely high fire weather and higher fuel loads. The areas of highest burning
1482 also showed an increase in human-driven burning. Further east (59 west, 7 south), the highest
1483 fire weather and decreased fuel moisture occurred alongside higher burning. Areas of higher
1484 BA in the north (63 west, 0 north) were associated with extreme fire weather, though offset by
1485 below average fuel loads. This may explain why they were not as extreme as the fires around
1486 Manaus, though the area is not generally considered fuel-limited (Kelley et al. 2019), and fuel
1487 had little overall impact across the region **Figure 10**).

1488 4 Attribution to Global Change Factors

1489

1490 4.1 Methods

1491

1492 Many of the direct drivers and controls on fire events, outlined in the **Section 3** (e.g. weather,
1493 fuel, moisture, ignition and suppression), are influenced by global change factors such as
1494 climate and land-use change. Since the pre-industrial era, global mean temperature has
1495 increased by between 1.1-1.3°C (Betts et al., 2023; Forster et al., 2024), with greater rates of
1496 warming at higher latitudes, adding potential for fuel drying. Climate change has also resulted
1497 in altered precipitation patterns, with total rainfall and dry season length increasing or
1498 decreasing variably across regions (Polade et al., 2014; Swain et al., 2018; IPCC, 2023a).
1499 Meanwhile, changes to fuel load and ignition rates are driven by climate change and
1500 anthropogenic land-use, with varying effects regionally (Finney et al., 2018; Romps, 2019).
1501 For example, in fuel-limited savannah biomes land-use change can drive more fragmented
1502 fuel loads and a reduction in fire (or an increase in fire resulting from land abandonment),
1503 whereas in forest ecosystems fragmentation provides more potential for ignition and leads to
1504 increases in fire occurrence (Andela et al., 2017; Rosan et al., 2022).

1505

1506 4.1.1 Overview of Attribution Approaches

1507

1508 In this report, we apply various modelling techniques for each focal region to attribute (i)
1509 regional changes in the probability of high fire weather to anthropogenic forcing (**Section**
1510 **4.1.2**) and (ii) changes in monthly BA to total climate forcing, socio-economic change, and all
1511 forcing (**Section 4.1.3**).

1512

1513 The types of forcing considered vary across the attribution techniques applied, and so we
1514 define here the terminology used throughout the paper when describing attribution results
1515 (summarised in **Table 4**).

1516

1517 Our attribution to *anthropogenic forcing* explicitly targets the changes driven by anthropogenic
1518 greenhouse gas emissions and land-use change, following the IPCC WGI definition (Hegerl
1519 et al., 2009; Mengel et al., 2021). We prescribe these emissions in a model to specifically
1520 isolate human forcing from natural variability (**Section 4.1.2**).

1521

1522 Our attribution to *total climate forcing* considers changes driven by climate change since the
 1523 pre-industrial period, including both anthropogenic forcing and natural variability in line with
 1524 the IPCC WGII and the Intersectoral Impacts Model Intercomparison Project 3a (ISIMIP3a)
 1525 definition of climate change impact attribution (IPCC, 2023b; IPCC 2023c; Mengel et al.,
 1526 2021). This involves comparing simulations driven with historical reanalysis to a detrended
 1527 counterfactual simulation with the historical warming signal removed (with both simulations
 1528 including historical transient land-use change) and therefore only the impacts of climate
 1529 change are attributed, not distinguishing between anthropogenic or natural causes (Mengel et
 1530 al., 2021; Burton, Lampe et al., 2023).

1531
 1532 Our attribution to *socio-economic factors* is applied via the same set of simulations as our
 1533 attribution to *total climate forcing*. The role of socio-economic factors is isolated by comparing
 1534 the early industrial period to the late industrial period in the counterfactual scenario, in which
 1535 only land-use and population density are allowed to change (Burton, Lampe et al., 2023).

1536
 1537 Finally, attribution to *all forcing* compares the early industrial period in the counterfactual
 1538 scenario to the last industrial period in the factual scenario, which gives the net effect of all
 1539 forcings combined. These are summarised in the **Table 4** below.

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Table 4: Summary of the attribution approaches used in this report.

Term	Definition	Experiments compared	Framework	Application
Event attribution for fire weather				
Anthropogenic Forcing	Change in fire weather driven by anthropogenic emissions from greenhouse gases, land-use change and aerosols. As per (Ciavarella et al., 2018; Li et al., 2021)	ALL: natural forcing plus human emissions NAT: Natural-only forcing from solar variability and volcanoes	HadGEM3-A attribution ensemble. 0.5 degree resolution	Fire weather (FWI)
Impacts attribution for burned area				
Total climate forcing	Changes in BA due to climate change, irrespective of the cause of warming. As per ISIMIP (Intersectoral Impacts Model Intercomparison Project) (Mengel et al., 2021 and Frieler et al., 2024)	Factual (2003-2019): present-day climate (driven by GSWP3-W5E5 reanalysis), CO ₂ , land-use and population Counterfactual (2003-2019): De-trended historical climate (warming signal removed), CO ₂ fixed at 1901 value, present-day land-use and population	ISIMIP3a impact attribution. 0.5 degree resolution	FireMIP ensemble and ConFire
Socio-economic factors	Changes in BA due to land-use and population change. As per (Burton, Lampe et al., 2023)	Counterfactual (1901-1917): Warming trend removed, fixed 1901 CO ₂ , limited land use and population change Counterfactual (2003-2019): Warming trend removed, fixed 1901 CO ₂ , present-day land use and population	ISIMIP3a impact attribution	FireMIP ensemble and ConFire
All forcing	Changes in BA due to climate, land-use and population change. As per (Burton, Lampe et al., 2023)	Counterfactual (1901-1917): Warming trend removed, fixed 1901 CO ₂ , limited land use and population change Factual (2003-2019): Historical climate driven by reanalysis	ISIMIP3a impact attribution	FireMIP ensemble

1543 The tools described here enable us to assess the influence of climate and socio-economic
1544 forcing on fire with respect to 3 different target variables. We use the FWI to assess how the
1545 probability of high (90th percentile) fire danger has changed as a result of anthropogenic
1546 forcing. As climate change has a direct impact on fire weather, this approach enables us to
1547 isolate its effects without confounding factors of land-use change and ignitions, and reveals
1548 how a fire might develop once ignited.

1549 In a second branch of analyses, we attribute the change in BA to total climate forcing, socio-
1550 economic factors, and all forcing, specifically targeting the observed month of peak burning in
1551 the 2023-24 fire season for each focal event. We use ConFIRE to assess the change in
1552 likelihood of a BA fraction (BA divided by the total area available to burn) that lies in the 90th
1553 or 95th percentiles of observations from the 2023-24 fire season (percentile thresholds vary
1554 regionally due to differing domain sizes).

1555 Separately, we attribute change in the monthly median BA in the present-day using
1556 simulations from fire-enabled dynamic global vegetation models (DGVMs) contributing to the
1557 Fire Model Intercomparison Project (FireMIP). Each of these methods is described in more
1558 detail below.

1559 In each approach we include an explicit estimate of uncertainty. We use bootstrapping to give
1560 uncertainty estimates for the FWI risk ratios. ConFire is designed as an uncertainty
1561 quantification model (see **Section 3.2.4**), giving the likelihood of all possible burned areas for
1562 each region based on a probabilistic analysis of past burn patterns and environmental conditions.
1563 We combine the information from the FireMIP models in a weighted multi-model ensemble to
1564 give uncertainty ranges across the models. Each result therefore presents a 5-95th percentile
1565 probability estimate.

1566 **4.1.2 Attributing Change in Likelihood of Fire Weather to Anthropogenic Forcing**

1567 We use an established approach to attribute change in probability of high (90th percentile) fire
1568 weather conditions to anthropogenic forcing. The approach uses estimates of FWI, as used in
1569 previous studies from the World Weather Attribution (Barnes et al., 2023), using outputs from
1570 the HadGEM3-A large ensemble (Christidis et al., 2013; Ciavarella et al., 2018). It follows the
1571 approach introduced by Stott et al. (2004) for attributing extreme weather events, and it has
1572 been employed in other attribution studies targeting fire weather, such as Li et al. (2021).
1573

1574 As outlined in **Section 3.1.1**, the FWI is used operationally and in research contexts to rate
1575 fire danger based on meteorological conditions. Due to the availability of model output
1576 variables we use maximum daily temperature at 1.5 m as a proxy for noon values, total daily
1577 precipitation, mean daily relative humidity at 1.5 m, and mean daily wind speed at 10 m,
1578 following Perry et al. (2022). We calculate the daily FWI for the month of 2023-24 peak BA
1579 anomaly for each focus region, using the same month and region for validation over the
1580 historical timeseries (1960-2013).
1581

1582 We validate and bias-adjust the model estimates of high FWI for the period 1960-2013 by
1583 comparing a 15-member HadGEM3-A ensemble with ERA5 reanalysis data (C3S, 2024)
1584 representing “observed” FWI. The 0.25 degree resolution observed FWI from ERA5 was
1585 coarsened by linear interpolation (calculated by extending the gradient of the closest two
1586 points) to match the 0.5 degree model grid. We compare the timeseries of individual
1587 components of the FWI (**Figure S40**), and the distribution of the modelled and observed FWI
1588 (**Supplementary Figures S41-S43**), and apply a simple linear regression to find the bias
1589 correction required for the 2023 model output. Before bias-adjustment, the modelled FWI is
1590 generally higher than the observed FWI, and some regions (e.g. Greece) require a larger
1591 correction than others. The correction adjusts the trend and absolute value while maintaining

1592 variability, and the model successfully reproduces the observed distribution after applying the
1593 correction in each region (see **Extended Methods and Evaluation Supplement**).

1594

1595 For the events occurring in the 2023-24 fire season, we calculate the FWI from the HadGEM3-
1596 A model simulations comprising 2 experiments of 525 members each, one driven by all
1597 forcings including historical greenhouse gas emissions, aerosols, zonal-mean ozone
1598 concentrations, land-use change and natural forcing (ALL), and a second counterfactual
1599 simulation with natural-only forcing from solar variability and volcanic emissions, and 1850
1600 land-use (NAT) (see **Table 4**). By applying the bias-adjustment from the previous step, and
1601 comparing the fire weather in the two simulations to the 2023 observed FWI from ERA5, we
1602 calculate the change in probability of high (90th percentile) fire weather due to anthropogenic
1603 forcing.

1604 **4.1.3 Attributing Change in Regional Burned Area to Total Climate Forcing, Socio-** 1605 **economic Factors and All Forcing**

1606

1607 **4.1.3.1 Peaks in Burned Area during 2023-24**

1608

1609 We use the ConFire attribution framework to attribute anomalies in BA fraction in the month
1610 of peak burning during the 2023-24 fire season to total climate forcing and socio-economic
1611 factors using the the Inter-Sectoral Impact Model Intercomparison Project (ISIMIP) 3a
1612 attribution protocol (see **Table 4**). For Canada and Western Amazonia, the attribution
1613 approaches are applied to cells with BA fractions in the 95th percentile of the BA fraction
1614 distribution during 2023-24. For Greece, the attribution approaches are applied to cells with
1615 BA fractions in the 90th percentile of the BA fraction distribution during 2023-24, with the lower
1616 percentile threshold selected due to the smaller domain size of Greece.

1617

1618 We trained ConFire on observed monthly BA from the MODIS BA product during 2003-2019
1619 at 0.5° across the entire region. For model training, we drive ConFire with Global Soil Wetness
1620 Project Phase 3 (GSWP3-W5E5) forcings provided at a 0.5° spatial resolution by ISIMIP3a
1621 (**Table 5**). The land surface information (tree cover and non-tree vegetated cover) is derived
1622 from the JULES-ES ISIMIP configuration (Mathison et al., 2023) driven by GSWP3-W5E5.
1623 This model includes dynamic vegetation, i.e changing vegetation cover in response to climate
1624 variables, growth, plant competition and mortality. So as not to double count the impact of fire,
1625 we turn the interactive vegetation-fire model off. The bias in this land surface information is
1626 adjusted to the MODIS Vegetation Continuous Fields collection 6.1 remote sensed data for
1627 <60°N (DiMiceli et al., 2022) and collection 6 for >60°N (DiMiceli et al., 2015) using a trend-
1628 preserving empirical quantile mapping bias adjustment method. This method significantly
1629 reduces the model bias in the JULES-ES output for most regions and variables, ensuring
1630 accurate means and distribution while preserving trends between historical and future periods
1631 (**Figure S22**). See “**Data and Data Processing**” in **Extended Methods Supplement** for
1632 details.

1633

1634 We ran ConFire in predictive mode on monthly timesteps with a structure similar to that used
1635 in **Section 3.1.2**, again grouping specific drivers into controls (**Table 5**). However, specific
1636 driving variables differed for this application: fuel load controls were represented by total
1637 vegetation cover and tree cover; fuel moisture controls were represented by mean consecutive
1638 dry days within each month, the fraction of dry days within the month, daily mean precipitation,
1639 mean and maximum monthly temperature, mean and maximum Vapour Pressure Deficit
1640 (VPD); ignition controls were represented by climatological lightning, pasture, crop and
1641 population density; and suppression controls were represented by pasture, crop and
1642 population density. ConFire’s Bayesian inference procedure proved useful because it allowed
1643 us to discern individual driver contributions (Gelman et al., 2013; Kelley et al., 2023). These

1644 modified ConFire simulations applied to all data in each of three experiments (see “simulation
1645 framework”).

1646

1647 To determine the impact of total climate forcing and socioeconomic factors on the increased
1648 BA during our focal events, we conducted a paired sampling of monthly BA in the target
1649 months (see **Table 4**). As there is no climate influence in the Early Industrial simulation, we
1650 first adjusted the target event (a monthly regional BA value) to that expected without climate
1651 change. For this adjustment, we find the percentile of the observed BA in the factual and find
1652 the BA at the same percentile in the counterfactual. We used paired samples to account for
1653 the uncertainty in the underlying mechanisms relating our drivers to BA, which would co-vary
1654 between experiments as per Kelley et al. (2021). In total, we took 200 samples over the 17
1655 years of each simulation, resulting in 3400 pairs.

1656

1657 The likelihood was then simply determined by the number of ensemble members in the
1658 factual scenario that predicted greater BA than the counterfactual for total climate forcing, or
1659 the counterfactual predicting greater BA than the early industrial scenario for socioeconomic
1660 factors. The relative increase in BA extent is the BA in factual over counterfactual (total
1661 climate forcing) or counterfactual over early industrial (socioeconomic).

1662

1663 As per **Section 3.1.2**, we evaluated the model following Barbosa (2024). We separately train
1664 ConFire on 50% of the data between 2003-2011 and perform evaluation on years 2012-2019.
1665 Further details of the model fitting and validation can be found in supplement **Extended**
1666 **Methods and Evaluation Supplement**.

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Table 5: Explanatory variables used for attributing extreme BA (**Section 4.1.3**) and multi-decadal outlook (**Section 5.1.2**). The explanatory variables are forcing data from the Inter-Sectoral Impact Model Intercomparison Project (ISIMIP) protocols ISIMIP3a and ISIMIP3b (Frieler et al., 2024). Positive (+ive) or negative (-ive) under “Controls” describes if a driver increases or decreases BA.

Variable	Controls	Construction	Source	Reference
Max. consecutive dry days	Moisture +ive	Monthly max of running count of days since rainfall > 0.1mm/m	Based on precipitation from ISIMIP3a/3b	Frieler et al. (2024)
No. dry days	Moisture +ive	fractional no. days of rainfall < 0.1mm/m	Based on precipitation from ISIMIP3a/3b	Frieler et al. (2024)
Maximum monthly temperature	Moisture +ive	maximum of maximum daily temperature within the month	Daily temperature approximated as ISIMIP3a/3b daily mean temperature + 0.5xdaily temperature range	Frieler et al. (2024)
Mean monthly temperature	Moisture +ive		ISIMIP3a/3b	Frieler et al. (2024)
Mean monthly Vapour Pressure Deficit (VPD)	Moisture +ive	mean of daily values constructed from specific humidity, surface pressure and max. temperature	ISIMIP3a/3b	Frieler et al. (2024); Barbosa (2024)
Maximum monthly VPD	Moisture +ive	max of daily values	ISIMIP3a/3b	Frieler et al. (2024); Barbosa (2024)
Tree Cover	Moisture -ive & Fuel +ive	JULES-ISIMIP annual mean tree cover bias-corrected to VCF vs JULES-ISIMIP3a factual interpolated to monthly from annual values	Joint UK Land Environment Simulator Earth System model (JULES-ES)-ISIMIP corrected to MODIS Vegetation Continuous Fields (VCF)	DiMiceli et al. (2017); Adzhar et al. (2022); Mathison et al. (2023)
Total vegetation cover	Fuel +ive	Tree cover plus non-tree vegetated cover simulated by JULES and bias-corrected as above	JULES-ES-ISIMIP corrected to MODIS VCF	DiMiceli et al. (2017); Adzhar et al. (2022); Mathison et al. (2023)
Lightning	Ignitions +ive	Climatology	LIS taken from ISIMIP3a	Kelley et al. (2014); Frieler et al. (2024)
Cropland	Ignitions +ive/ Suppression -ive	Interpolated from annual to monthly	ISIMIP3a/3b	Frieler et al. (2024))
Pasture	Ignitions +ive/ Suppression -ive	Interpolated from annual to monthly	ISIMIP3a/3b	Frieler et al. (2024)
Population Density	Ignitions +ive/ Suppression -ive	Interpolated from annual to monthly	ISIMIP3a/3b	Frieler et al. (2024)

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4.1.3.2 Background Changes in Burned Area during 2003-2019

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Finally, we attribute changes in median monthly BA across all months in 2003-2019 to total climate forcing, socio-economic factors and all forcings, using the novel attribution method developed using state-of-the-art global fire models from the FireMIP (Burton, Lampe et al., 2023). This represents an assessment of how BA has changed during the 2003-2019 period versus counterfactual scenarios. Our method employs the same ISIMIP3a simulation framework outlined above with 7 fire-enabled DGVMs (see **Table S1**) for the period 1901-2019 for the factual and counterfactual experiments (see **Table 4** for descriptions). The

1684 ConFire was not used in this element of our attribution approaches; rather, the native fire
1685 modelling scheme of each fire-enabled DGVM was employed. Model fire schemes are
1686 described in Burton, Lampe et al. (2023).

1687 A weighted ensemble of the monthly outputs of BA, based on the regional performance of the
1688 unweighted models against observational data from GFED5 and FireCCI, is used for the
1689 analysis. Due to large differences in absolute values of BA between the GFED5 and FireCCI
1690 observational datasets and across the models, the weightings in the ensemble are based on
1691 model capability to capture relative anomalies present in the observational datasets on a
1692 regional basis, and all changes are reported as relative anomalies. We focus on the change
1693 in median monthly BA across all months in 2003-2019 because the fire models underpredict
1694 the high tails of the distribution. The weighted models are randomly resampled to generate
1695 uncertainty estimates for each region. The method and results are reported in full for all 43
1696 IPCC AR6 regions in Burton, Lampe et al. (2023), and in the current report we select the IPCC
1697 regions that align most closely with our focus regions defined in **Section 3.2**.
1698

1699 **4.2 Results**

1700 **4.2.1 Change in the Likelihood of High Fire Weather in 2023-24**

1703 **4.2.1.1 Canada**

1704
1705 The fire weather conditions in Canada during June 2023 were 2.9-3.6 times more likely due
1706 to anthropogenic forcing. Here we assess the 95th percentile of FWI over the country during
1707 the month of peak anomaly in BA (June) in the ALL and NAT HadGEM3 simulations. More of
1708 the ALL distribution lies above the observed 95th percentile of FWI from ERA5 compared to
1709 the NAT distribution (**Figure 13**), and we therefore conclude that the probability of
1710 experiencing the high fire weather observed during June 2023 is more likely in a climate forced
1711 with anthropogenic emissions.
1712

1713 **4.2.1.2 Greece**

1714
1715 The high fire weather conditions experienced during the peak anomaly in BA in August 2023
1716 were 1.9-4.1 times more likely due to anthropogenic forcing (**Figure 13**). In this case the 95th
1717 percentile of FWI is outside of the distribution so we instead assess the 90th percentile of FWI
1718 over the country. This is likely a result of our linear inference of 2023 for the bias correction
1719 based on the 1960-2013 period, where in fact 2023 was so anomalous that it doesn't fit this
1720 trend. The 2023 event threshold here also lies at the very high end of simulated fire weather,
1721 meaning it was very unusual in the model simulations. The result range here is also larger
1722 than for Canada, meaning there is less certainty about how much human influence has
1723 increased the probability, although it does highlight at least a 50% increase in likelihood of
1724 high fire weather.
1725

1726 **4.2.1.3 Western Amazonia**

1727
1728 High fire weather in western Amazonia during Sept-Oct 2023 was 20.0-28.5 times more likely
1729 due to anthropogenic forcing (**Figure 13**). In this region there is a large shift in the ALL forcing
1730 distribution compared to the NAT only forcing for the 95th percentile of FWI, and the high risk
1731 ratio shows a strong anthropogenic signal in driving the meteorological conditions that led to
1732 high fires over this period.
1733

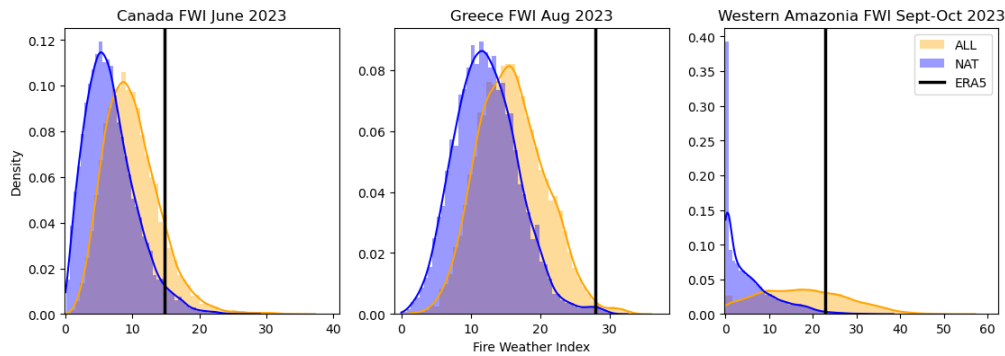


Figure 13: High FWI in 2023: **(left)** 95th percentile FWI in June over Canada, **(middle)** 90th percentile FWI in August over Greece and **(right)** 95th percentile FWI in September-October over western Amazonia, in the HadGEM3 ensemble of ALL (anthropogenic plus natural forcing, orange) and NAT (natural-only forcing, blue) bias-adjusted simulations, and ERA5 reanalysis (black line).

4.2.2 Change in the Likelihood of Peaks in Burned Area in 2023-24

4.2.2.1 Canada

We show that total climate forcing was virtually certain (99.9% confidence) to have led to greater BA in Canada during June 2023 (**Figure 14**). Our attribution results indicate that total climate forcing increased BA extent by 4.1 [2.3-7.8]% during the month of June across the period 2003-2019. Additionally, considering the anomalies in fire drivers during 2023, we estimate an additional 5.7 [1-30]% absolute increase in BA extent during June 2023, on top of the 2003-2019 period (**Figure 10**). The impact of socio-economic factors is less certain, with only a 64.8% likelihood of decreasing burning, affecting BA extent by between -22.5 and 6.6% during 2003-2019.

Overall, we estimate that BA in Canada in June 2023 was 10.1 [3.3-40.1]% greater due to total climate forcing in the 2003-2019 period combined with this year's anomaly in the climatic variables. As a caveat, we note that this is not a formal attribution of the 2023 anomaly because no counterfactual exists for the year, but rather an attribution of the change in BA in 2003-2019 with the additional influence of climate factors on BA in 2023 superimposed (this caveat also applies to the other focal regions). For Canada, the BA attribution targets cells in the 95th percentile of the BA fraction distribution.

1763 **4.2.2.2 Greece**

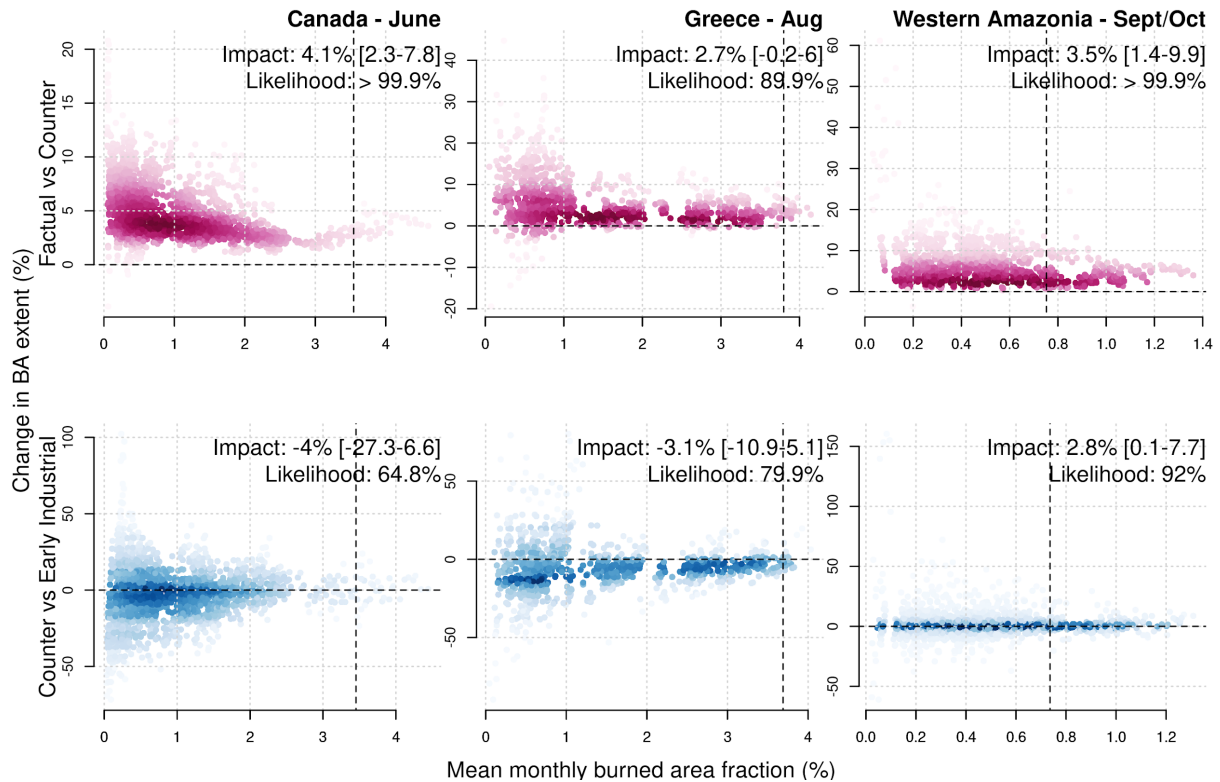
1764
1765 Total climate forcing caused a change in the likelihood of high BA in Greece of 2.7 [-0.2 to
1766 6.0]% in the period 2003-2019 in the cells with the greatest BA fraction (90th percentile), with
1767 90% confidence (**Figure 14**). In this case we use the 90th percentile to represent high BA over
1768 the region for the month of August, over 2003-2019. This increase is likely a conservative
1769 figure given the additional warming since 2019, and estimating the additional burning that
1770 might have been experienced during the anomalous conditions of 2023, we find an additional
1771 change of 1.3 [-9.3-11]% (**Figure 10**). Socioeconomic factors likely (79.9%) caused a
1772 decrease in burning, though could have caused an increase, affecting BA extent by -3.1 [-10.9
1773 to 5.1]%.
1774

1775 Overall we estimate that BA in Greece in August 2023 was increased by 4 [-9.5 to 17.7]% due
1776 to total climate forcing in the 2003-2019 period combined with this year's anomaly in the
1777 climatic variables. In the case of Greece, uncertainties around the influence of total climate
1778 forcing and socioeconomic factors are greater because the smaller region size limits
1779 information available for model optimization. For Greece, the BA attribution targets cells in the
1780 90th percentile of the BA fraction distribution.
1781

1782 **4.2.2.3 Western Amazonia**

1783
1784 Over the period 2003-2019, total climate forcing was virtually certain to have caused an
1785 increase in burned areas like the one experienced in western Amazonia in September and
1786 October 2023 (>99.9% likelihood), with a likely range of increase in extent of 3.5 [1.4-9.9]%
1787 (**Figure 14**). Here we assess the change in BA due to total climate forcing in the cells that
1788 burn most regularly (95th percentile) of our defined region of western Amazonia over Sept and
1789 Oct 2003-2019. Extending our analysis to the 2023 anomaly, we estimate that additional
1790 burning could have been up to 7.4 [0.8-36.2]% on top of the 2003-2019 levels. Despite finding
1791 little influence of humans specifically in 2023 compared to the previous 10 years in **Figure 10**,
1792 since early-industrial, we show socioeconomic factors have had a large influence on the
1793 occurrence of extreme levels of burning. Events similar to 2023 were very likely exacerbated
1794 by socioeconomic conditions (92.0% likelihood) increasing BA by 2.9 [0.1-7.8]%.
1795

1796 Overall we estimate that BA in western Amazonia in September-October 2023 was increased
1797 by 11.1 [2.2-49.7]% due to total climate forcing in the 2003-2019 period combined with this
1798 year's anomaly in the climatic variables. For western Amazonia, the BA attribution targets cells
1799 in the 95th percentile of the BA fraction distribution.
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1801



1802
 1803 **Figure 14:** Estimated change in BA in selected months in the period 2003-2019 (e.g.
 1804 seventeen Junes for Canada) as modelled by ConFire when driven by the ISIMIP3a reanalysis
 1805 versus counterfactual scenarios. The panels shows the change in BA extent (%) due to (**top**
 1806 **row**) total climate forcing versus a scenario without climate forcing and (**bottom row**)
 1807 socioeconomic forcing versus a scenario without socioeconomic forcing. Results are shown
 1808 for (**left**) Canada, (**middle**) Greece, and (**right**) Western Amazonia. Many points are plotted
 1809 because each point represents a posterior estimate of change in BA for a month and there
 1810 are 1,000 iterations of ConFire to explore the effect of uncertainty in input parameters and
 1811 structural relationships between BA and input variables. Impact values are the central
 1812 (median) and 5th-95th percentile range of the estimated values of change in BA (%). The
 1813 likelihoods shown are the probability of the change being significant, measured as the % of
 1814 the posterior estimates being above/below zero when the median is above/below zero.

1815 1816 1817 **4.2.3 Background Changes in Burned Area due to Total Climate Forcing, 1818 Socioeconomic Factors, and All Forcing**

1819 1820 **4.2.3.1 Canada**

1821
 1822 As reported in Burton, Lampe et al. (2023), we also show how background levels of fire extent,
 1823 represented by the modelled median BA for months of interest, has changed overall in Canada
 1824 due to total climate forcing (**Figure 15**), socioeconomic forcing (**Figure S15**), and all forcings
 1825 (**Figure S16**). Using AR6 regions that best match our focus areas, we show that BA has
 1826 increased by 1.9% [0.1, 3.6] in North West North America (NWN) due to total climate forcing,
 1827 but reduced by -0.2% [-1.7, 1.3] in North East North America (NEN) (**Figure 15**). In these
 1828 regions, socioeconomic forcing has dampened the effects of climate change, by reducing BA
 1829 by -9.5% [-13.6, -6.3] in NWN, and -8.5% [-12.5, -5.7] in NEN (see Supplement). All forcings
 1830 combined have led to an overall reduction in BA of -8.3% [-12.5, -4.9] in NWN and -8.7% [-
 1831 12.8, -5.8] in NEN (**Figure S16**).
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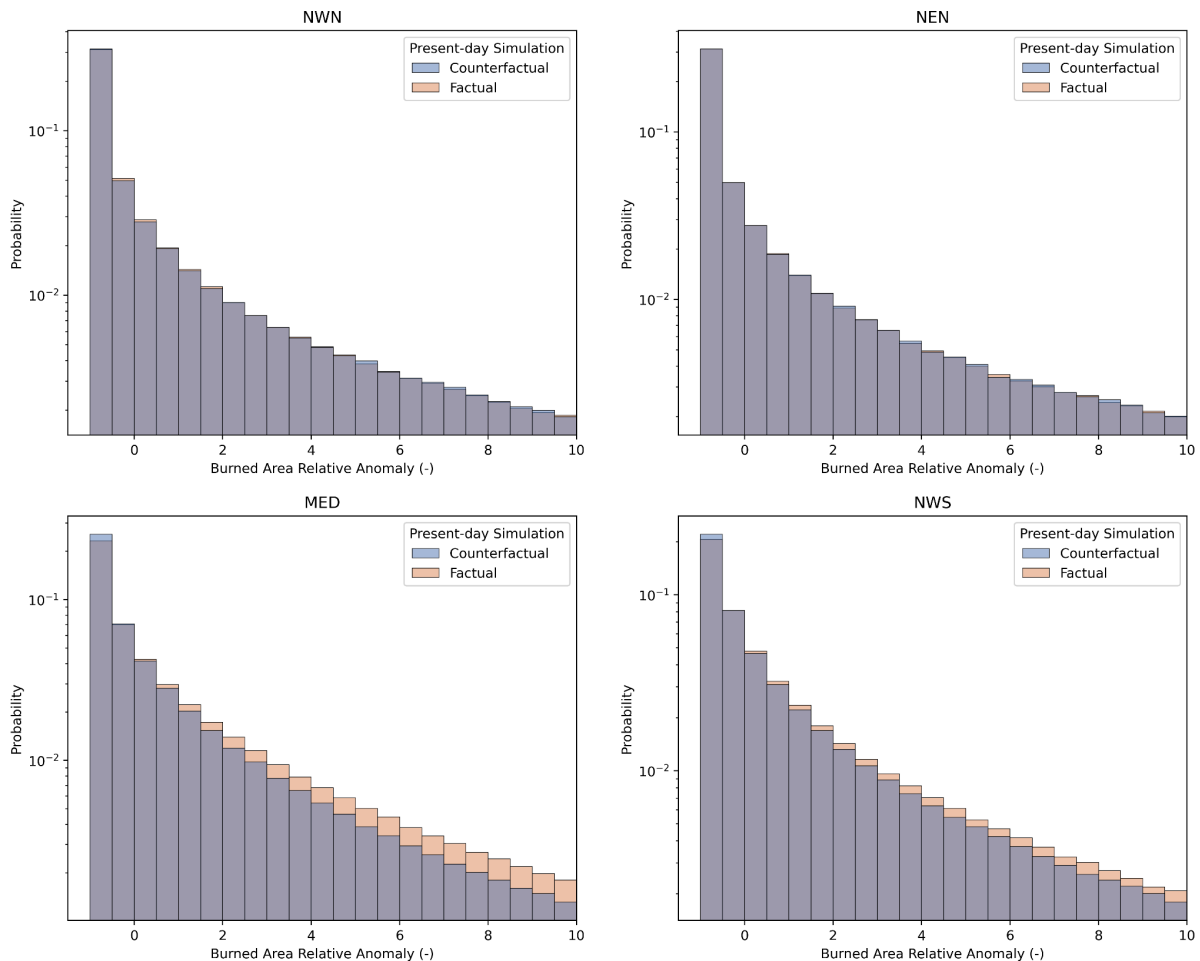
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4.2.3.2 Greece

Burton, Lampe et al. (2023) find a larger increase in median BA for months of interest due to total climate forcing in the Mediterranean region (MED), with an increase of 14.5% [11.5, 18.1] today compared to the counterfactual (**Figure 15**). This is particularly the case for the high burned areas, where the increase is larger compared to the lower end of the distribution. However, socioeconomic factors have largely offset this by reducing BA by -10.2% [-13.6, -6.6]. All forcings combined have led to an overall regional increase in BA of 0.5% [-3.5, 5.5] (see **Figure S16**).

4.2.3.3 Western Amazonia

As per Burton, Lampe et al. (2023), total climate forcing has increased median BA for months of interest by 11.5% [5.4, 18.4] in Northwest South America (NWS) today compared to the counterfactual (**Figure 15**). Again, this increase is mostly impacting the BA at the higher end of the distribution. This is mostly offset by socioeconomic factors (-9.0% [-18.9, 1.2]), although all forcings combined have still led to an overall increase in BA of 1.5% [-6.9, 10.5] in the region (see Supplement).



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Figure 15: Change in median BA due to total climate forcing from FireMIP. Present day BA (2003-2019) for factual (historical forcing, orange) and counterfactual (detrended climate, blue), for AR6 regions. Panels show (**top left**) North West North America (NWN), (**top right**) North East North America NEN, (**bottom left**) Mediterranean (MED), and (**bottom right**) North West South America (NWS). Probability is shown on a log scale.

1860 **5 Seasonal and Multi-Decadal Outlook**

1861

1862 **5.1 Methods**

1863

1864 **5.1.1 Seasonal Forecasts**

1865

1866 Among the modes of variability in the climate system most relevant to wildfire activity globally
1867 is the El Niño-Southern Oscillation (ENSO) (Mariani et al., 2016; Fuller and Murphy, 2006;
1868 Cardil et al., 2023). Numerous studies have demonstrated that there is a predictable cascade
1869 of fire across tropical continents during ENSO events, highlighting staggered responses of
1870 wildfire to ENSO (Chen et al., 2017). The utility of using ENSO as a predictor of fire is
1871 highlighted by its application in Indonesia, where severe fires during the 2015 El Niño led to
1872 hazardous haze in Singapore and diplomatic tensions in the region, prompting better regional
1873 cooperation and enforcement of anti-burning laws (Field et al., 2016; Forsyth, 2014; Carmenta
1874 et al., 2021). Consequently, Indonesia now implements preemptive bans on agricultural
1875 burning based solely on ENSO predictions, a measure that proved successful in 2023 when
1876 no significant fire anomalies were recorded despite a strong positive ENSO event (Lin et al.,
1877 2020; Sloan et al., 2022).

1878

1879 Another phenomenon demonstrably linked to global fire activity is the Indian Ocean Dipole
1880 (IOD), which occurs in the Indian Ocean. There is ongoing debate regarding the direct
1881 influence of the IOD on Australian fires, for example as the signal is often modulated by
1882 changes in land management practices (Harris and Lucas, 2019). Other atmospheric modes
1883 of variability in the Southern, Northern hemisphere and in the arctic regions can also have
1884 influence on interannual BA patterns, and **Figure S1** shows the climate modes with strongest
1885 influence on regional BA globally.

1886

1887 Outputs available from the Copernicus Climate Change Service (C3S) multi-model seasonal
1888 prediction system are used to evaluate large-scale climate modes with the most proven links
1889 to variation in fire activity: ENSO and IOD (CDS, 2018). As not all regions display similar
1890 seasonal direct correlations between fire activity and ENSO, we also use seasonal outlooks
1891 of the FWI from one of the models, ECMWF-SEAS5, to identify probabilities for the
1892 establishment of anomalous landscape flammability in the next season. This is done using a
1893 51-member forecast ensemble and a 24-year model climatological distribution (derived from
1894 a 25-member ensemble re-forecast) covering the period 1993-2016. The probability of
1895 exceedance is determined based on the proportion of forecast members meeting an
1896 distributional threshold at any given geographical point. We consider the 75th percentile
1897 threshold indicative of moderate anomalous conditions and the 95th percentile indicative of
1898 extreme anomalous conditions.

1899

1900 **5.1.2 Decadal Projections of Burned Area**

1901

1902 In order to project future changes in BA, we utilised the same modelling approach detailed in
1903 **Section 4.1.3.1**, following a similar protocol to UNEP (2022a). We drive the ConFire model
1904 with ISIMIP3a and bias-corrected JULES-ES data. For predictive mode, we used bias-
1905 corrected global climate model (GCM) outputs from ISIMIP3b. While ISIMIP3a provides
1906 reanalysis datasets to drive models for impact assessments, ISIMIP3b provides driving data
1907 from 5 bias-corrected GCMs, including historical data up to 2014 and future scenarios from
1908 2015-2100 under Shared Socioeconomic Pathway (SSP) scenarios SSP126, SSP370, and
1909 SSP585. Each SSP represents future socio-economic pathways and includes GHG emissions
1910 to drive the GCMs. The 5 GCMs used are: GFDL-ESM4 (Held et al., 2019), IPSL-CM6A-LR
1911 (Boucher et al., 2020), MPI-ESM1-2-HR (Mauritsen et al., 2019), MRI-ESM2-0 (Yukimoto et

1912 al., 2019), and UKESM1-0-LL (Tang et al., 2019; Sellar et al., 2019). As part of ISIMIP3b, each
1913 GCM is bias-corrected as described in Lange (2019).

1914 At present, future projections for land use and population density forcing were not available
1915 for ISIMIP3b, so we only considered the influence of climate and vegetation fuel loads (related
1916 to land cover) on fire, and not changes in ignitions or land use. We used JULES-ES land cover
1917 outputs as per the previous section, but this time with JULES driven by each of the 5 different
1918 bias-corrected GCMs and for the 3 different SSP scenarios instead of historical reanalysis.
1919 The land cover output was then bias-corrected (using the same mapping procedure as
1920 **Section 4.1.3.1**, based on biases between JULES-ES driven by reanalysis and VCF
1921 observations) to maintain consistency with the GCM bias-correction procedures. We apply an
1922 additional bias-correction to preserve the trend in vegetation cover from the historical period,
1923 and to smooth the transition between the historical and future periods (see “**Data and Data**
1924 **Processing**” in **Extended Methods Supplement** for details). The results in **Section 5.2.2**
1925 are for the months June-August for Canada, July-September or Greece and August-October
1926 for Western Amazonia, corresponding to those regions' fire seasons today.

1927 Our approach provides a probability distribution of future BA representing the uncertainty
1928 range from cross-model (GCM) spread in the response of climate and vegetation to emissions
1929 for each scenario and year in the period 2010-2100. Years 2010-2014 were consistently
1930 adopted from the historical experiment.

1931 For Western Amazonia focal event, we additionally tested a 1-in-100 year event under 2010-
1932 2020 climate, defined as the BA at the 99th percentile ConFires distribution. We also use the
1933 1-in-100 definition at a grid cell level to determine spatial variations in the change in extreme
1934 fire for each region. We then calculated decadal average likelihoods of the regions' event in
1935 each decade up to 2100. Return times are 1/likelihood. The change in likelihood of an event
1936 occurring on a given return time was calculated relative to the 2010-2020 baseline period.

1937 For the Canada focal event, we also calculated the integrated probability of an event with
1938 similar magnitude to 2023 within the expected lifespan of a Canadian citizen. According to UN
1939 population statistics, the average life expectancy of a Canadian citizen born today is 83 years
1940 (United Nations Population Division, 2022). In order to cover the 7-year period after 2100, we
1941 extrapolated the annual trend in probabilities. The integrated probability is calculated as one
1942 minus the product of the annual probability of not seeing a fire event like 2023, for each year
1943 between 2023 and 2106.

1944

1945 **5.2 Results**

1946

1947 **5.2.1 Seasonal Outlook for 2024**

1948

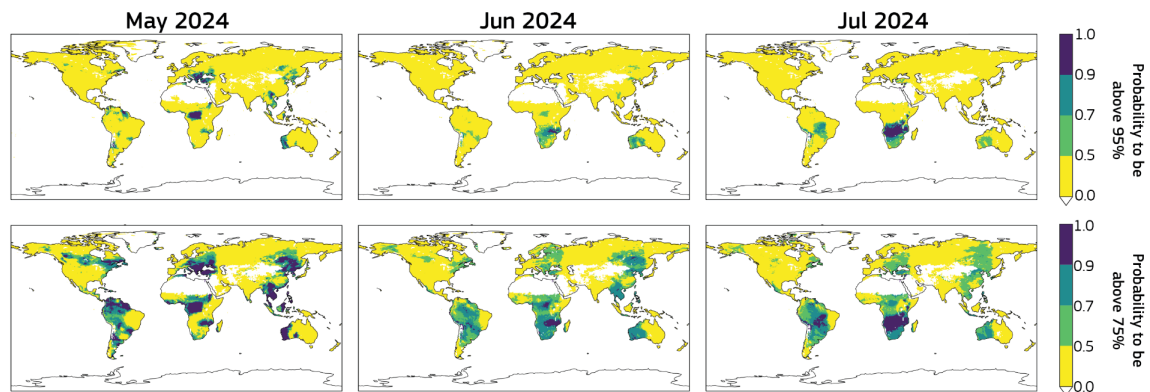
1949 The 2023–2024 El Niño event emerged as the fourth-most powerful on record, causing
1950 widespread droughts, floods and other anomalous conditions worldwide. Officially declared by
1951 the World Meteorological Organization (WMO) on July 4, 2023, its meteorological impacts
1952 have unfolded between November 2023 and April 2024 (Joshi, 2023). Climate scientists have
1953 found that the 2023–24 El Niño event, superimposed on climate change signal, has elevated
1954 global temperatures beyond the records set during the 2016 El Niño event. Global mean
1955 surface temperatures in 2023 were 1.31°C above pre-industrial levels of 1850-1900 (Forster
1956 et al., 2024).

1957

1958 As of mid-June 2024, El Niño conditions have transitioned toward neutral conditions, which
1959 are forecasted to persist during the boreal summer. The Indian Ocean Dipole (IOD) index is
1960 currently positive, and forecasts indicate that it will remain in a positive state for the next
1961 season. Connections are established between a positive IOD phase and fire risk in Indonesia

1962 and parts of Australia, though outcomes generally depend on interactions with the ENSO
 1963 phase (Pan et al., 2018; Ren et al., 2024; Abram et al., 2021). Similarly, the positive phase of
 1964 the IOD is linked with heightened fire risk in South America, particularly in the Amazon basin,
 1965 where it can interact with other climate teleconnections to exacerbate droughts (Cardil et al.,
 1966 2023; Dong et al., 2021).

1967
 1968 **Figure 16** shows the predicted probabilities of the monthly average FWI exceeding moderate
 1969 (75th percentile) or high (95th percentile) thresholds of the monthly climatology. Most areas in
 1970 Southeast Asia and South America are predicted to experience a decrease in the likelihood of
 1971 anomalous conditions over May, June and July 2024. Parts of Canada are predicted to reach
 1972 moderate anomalous conditions once again in early summer, and this combined with
 1973 overwintering fires could promote a second consecutive high fire season as has already been
 1974 reported in the media (*BBC News*, 2024, *New York Times*, 2024). Predictions also suggest
 1975 that the moderate FWI threshold will be exceeded in southeast, central, and western Brazil
 1976 with the high threshold exceeded in southern and western parts. In parts of Africa, moderate
 1977 FWI anomalies may be experienced throughout June-August.
 1978



1979
 1980 **Figure 16:** Probability for the monthly average FWI exceeding the **(top row)** 95th percentile
 1981 threshold (anomalous conditions) and **(bottom row)** 75th percentile threshold (extremely
 1982 anomalous conditions) of the monthly climatological distributions. The probability is calculated
 1983 using the 51 ensemble member realisation from ECMWF’s long-range forecasting system,
 1984 ECMWF-SEAS5 FWI, and comparing them with the 1991-2016 climatology (Copernicus
 1985 Emergency Management Service, 2019).

1986 5.2.2 Future Changes in Likelihood of Extreme Fire Events

1987 5.2.2.1 Canada

1988
 1989
 1990
 1991 The probability of Canada experiencing BA extent similar to June 2023 (specifically, for cells
 1992 with BA fraction in the 95th percentile in that month) is estimated to be 0.15% in any given year
 1993 under the climate conditions of 2010-2020, according to estimates made using reanalysis data
 1994 (**Table 6; Figure 17**). Bias-correction did not fully resolve all discrepancies between the GCMs
 1995 and reanalysis data, and the GCMs gave a range of likelihoods spanning 0.02 to 0.71% for
 1996 any given year under the climate conditions of 2010-2020. We describe future changes as
 1997 significant if the range across GCM projections for a future period does not overlap with the
 1998 range given by the GCMs for 2010-2020.
 1999

2000 By the 2040s, the likelihood of an event like 2023 increases significantly to 0.42% - 2.2%
 2001 across scenarios, which is two to six times as likely as in the 2010s (**Table 6; Figure 17**).
 2002 While the likelihood of an event like 2023 occurring in the 2040s is slightly higher in SSP585

2003 (0.6%-2.2%) than other scenarios (0.41%-1.6% for SSP126 and 0.5%-1.7% for SSP370),
2004 differences between future scenarios are not significant at the mid-century point.
2005

2006 The SSP126 scenario diverges significantly from SSP370 and SSP585 after 2070. Under
2007 SSP126, the likelihood of an event like 2023 stabilises at 0.3-0.8% in the 2070s and remains
2008 largely unchanged until the 2090s (**Figure 17**). In contrast, the likelihood of an event like 2023
2009 continues to rise to 2.1-3.7% in the 2090s under SSP585 (**Table 6, Figure 17**). Under
2010 SSP126, the probability of at least one event like 2023 recurring in any year (of any decade)
2011 between 2024 and 2100 is estimated to be 18%-73%, compared with 59-87% under SSP585.
2012 Hence the probability of an event like 2023 recurring by the 2090s is estimated to be twice
2013 greater in SSP585 than in SSP126.
2014

2015 Someone born in Canada in the current decade, with a life expectancy of 83 years (United
2016 Nations Population Division, 2022), has a 65-90% probability of seeing a similar event in their
2017 lifetimes under SSP585, compared with only an 12% likelihood of someone who reached 83
2018 years old in the 2010s. Someone born in Canada today would also have a 42-80% probability
2019 of seeing an event of similar magnitude *twice* under SSP585. Under SSP370, a citizen has a
2020 48-84% probability of seeing a similar event once and 29-76% probability of seeing a similar
2021 event twice. This reduces to 19-76% for one occurrence and 4-58% for two occurrences under
2022 SSP126.
2023

2024 These differences outlined above highlight the divergence in future likelihood of a 2023-like
2025 event in Canada between high-mitigation (SSP126) and no-mitigation (SSP585) scenarios.
2026 The divergence of likelihoods between the two scenarios is associated with increases in both
2027 fuel load and fuel dryness (**Figure 17**).
2028

2029 Future changes in probability of an event like 2030 under the SSP370 lie closer to those of
2030 SSP585 than SSP126 (**Table 6; Figure 17**). For example, the probability of an event like 2023
2031 occurring at least once by the 2090s is estimated to be between 48% and 84% under SSP370,
2032 only slightly below the range projected under SSP585. This highlights that high-mitigation, low
2033 emissions pathways are required to limit future increases in the potential for high-impact
2034 events in Canada this century.
2035

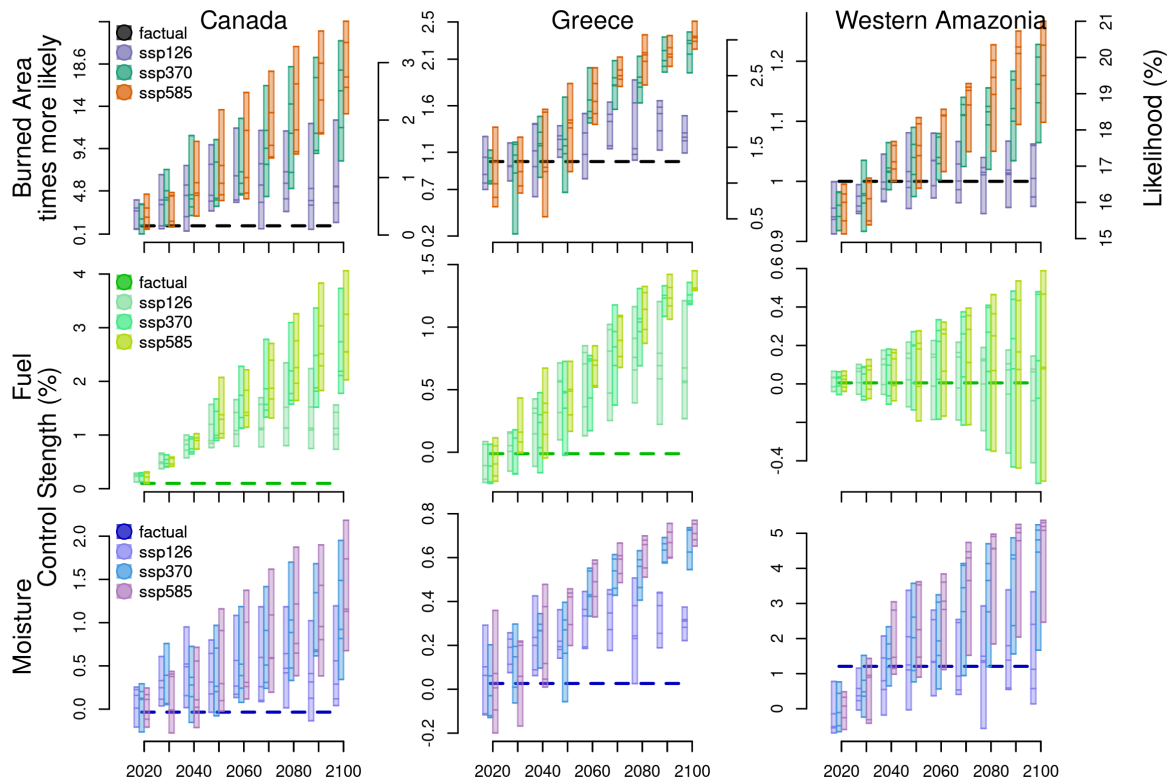
2036 **Figure 18** shows the spatial variability of changes in summer BA and 1-in-100 year BA events
2037 in Canada. While some areas see increases in BA and fire extremes in all scenarios, the
2038 greatest rates of change are projected in Southern Alberta and Saskatchewan and under
2039 SSP585. These patterns emerge as early as 2030 (**Figure S17**). Meanwhile, the Yukon and
2040 Northwest Territories are projected to see increased BA from 2040 in all scenarios, with 1-in-
2041 100 year BA events becoming around twice as likely in parts of these provinces (**Figure S18**).
2042 By the end of the century, a larger increase in BA is seen under SSP585 and SSP370 than in
2043 SSP126, with 1-in-100 year BA events becoming up to 5 times more likely in parts of Yukon
2044 and Northwest Territories. Uniquely under SSP585, factor-two increases in BA and 1-in-100
2045 year BA events extend into British Columbia.

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Table 6: Summary of the likelihood of extreme events today using reanalysis ‘factual’ and today and into the future using bias-corrected GCMs for our three focal regions. ‘2023’ events focuses on the BA extreme identified in **Section 3.4.3**. 1-in-100 for Western Amazonia additionally looks at the likelihood of a 1-in-100 event under present day climate conditions, following the definition of extreme in (UNEP, 2022a). We also determine how much more frequent the events will be at two different time horizons based on each models likelihood in the future projections over likelihood during 2010-2020. Asterisks (*) indicate non-significant changes from 2010-2020 values. Colours show linear increase of likelihood (red) and frequency (orange), where darker shade indicates higher values.

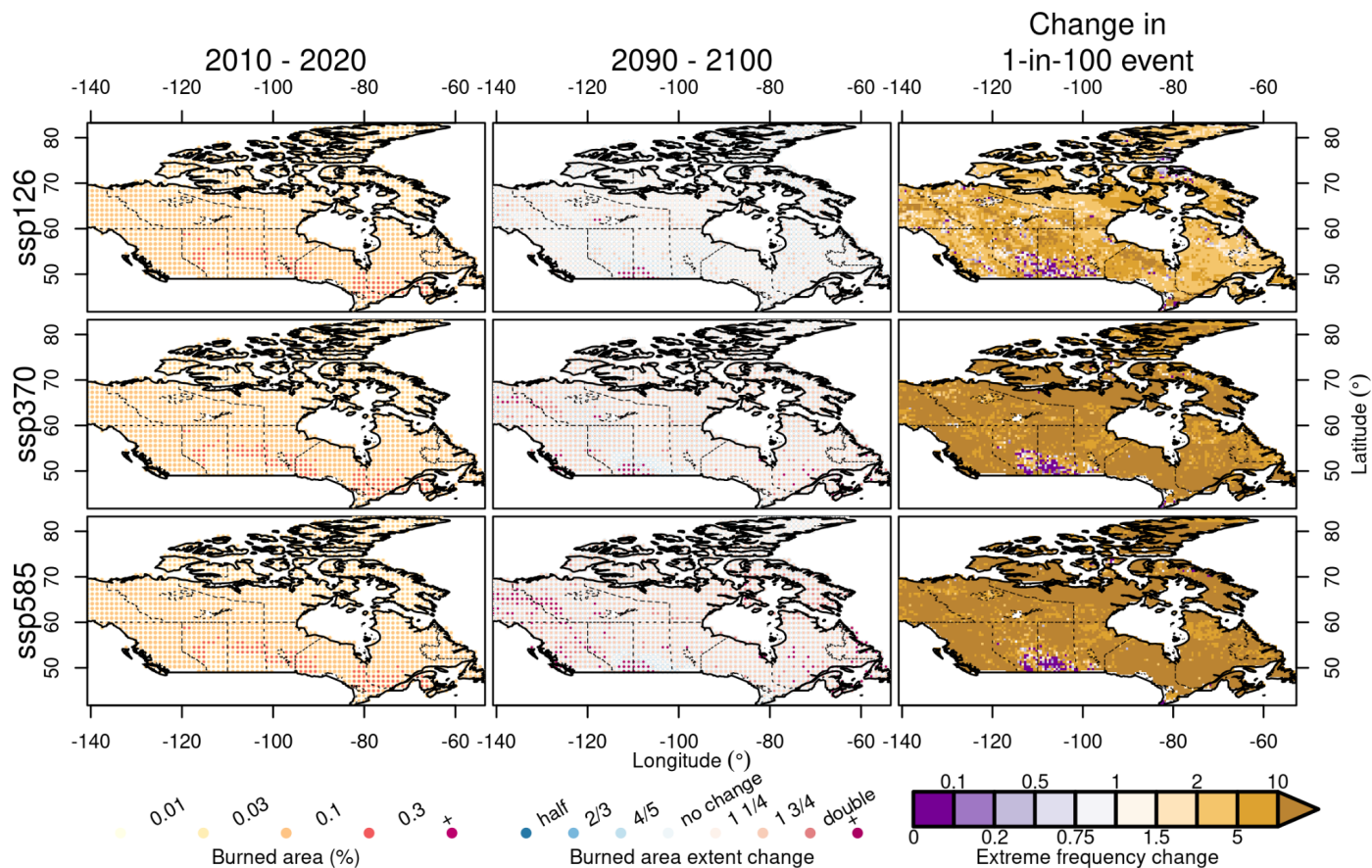
Region	Event	SSP	Represents	Likelihood (%/year)						How much more frequent (multiplier)			
				2010-2020		2040-2050		2090-2100		2040-2050		2090-2100	
				<i>min</i>	<i>max</i>	<i>min</i>	<i>max</i>	<i>min</i>	<i>max</i>	<i>min</i>	<i>max</i>	<i>min</i>	<i>max</i>
Canada	2023 (~1-in-700)	Factual	observed	0.15									
		SSP126	strong mitigation	0.1	0.61	0.42*	1.6*	0.22*	2*	2.6*	4.2*	2.2*	3.3*
		SSP370	middle of the road	0.12	0.54	0.5*	1.7*	1.3	3.4	3.1*	4.2*	6.3	10.8
		SSP585	no mitigation	0.1	0.71	0.6*	2.2*	2.1	3.7	3.1*	6*	5.2	21.1
Greece	2023 (~1-in-80)	Factual	observed	1.3									
		SSP126	strong mitigation	0.91	1.7	1.4*	1.8*	1.4*	1.9*	1.1*	1.5*	1.2*	1.6*
		SSP370	middle of the road	0.99	1.5	0.87*	2.2*	2.5	3.1	0.9*	1.5*	2.1	2.6
		SSP585	no mitigation	0.67	1.8	1.2*	2.4*	2.9	3.3	1.3*	1.7*	1.8	4.3
Western Amazonia	2023 (~1-in-6)	Factual	observed	16.6									
		SSP126	strong mitigation	15.1	16.6	15.8*	17.9*	15.9*	17.6*	1.1*	1.1*	1.1*	1.1*
		SSP370	middle of the road	15.2	16.3	16.2*	18.1*	17.7	20.4	1.1*	1.1*	1.2	1.3
		SSP585	no mitigation	15.1	16.5	16.4*	18.3*	18.2	21	1.1*	1.1*	1.2	1.3
	1-in-100	Factual	observed	1.5									
		SSP126	strong mitigation	0.82	1.5	1.2*	2.2*	1.2*	2*	1.4*	1.5*	1.3*	1.5*
		SSP370	middle of the road	0.81	1.5	1.3*	2.2*	2	3.1	1.5*	1.6*	2.2	2.4
		SSP585	no mitigation	0.8	1.5	1.4*	2.4*	2.3	3.3	1.6*	1.8*	2.2	2.9

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Figure 17: Future projections from ConFire of the likelihood of BA extremes of the magnitude seen in 2023-2024, and their corresponding fuel and moisture controls. Each set of bars shows changes in each decade, with each bar representing a different SSP scenario and the spread of bars indicating the variation across GCMs, with individual bars representing different GCMs. The BA data indicates (**right axis**) the likelihood of the defined event occurring in any given year for future decades and (**left axis**) the frequency of occurrence relative to reanalysis-driven simulations (factual ISIMIP3a – dashed line) of the years 2010-2020. The fuel and moisture rows illustrate the relative % change in the maximum burning allowed by that control compared with reanalysis-driven simulations for the years 2010-2020.



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Figure 18: Projected changes in June-August BA over Canada by 2090-2100 under three SSP scenarios, with BA simulated by ConFire. **(Left)** Average June-August BA fraction (%) for 2010-2020. **(Middle)** Relative change in June-August BA extent projected for 2090-2100 period, expressed as a multiplier of 2010-2020 values. **(Right)** Increased (or decreased) frequency in the 2090-2100 period of a 1-in-100 year event defined for the period 2010-2020, expressed as a multiplier of 2010-2020 values. In the left column, the size of the dot in each grid cell indicates the likelihood (larger = higher likelihood) of a BA fraction (or being greater than the threshold indicated by the coloured dot (see legend at the base)). Likewise, the size of the dot varies with likelihood that the BA fraction exceeds the threshold indicated by the coloured dot (see legend at the base). For example, a large pale orange dot in the left column indicates a high likelihood of the BA fraction exceeding 0.05%, whereas a small dark red dot indicates a small (but non-zero) likelihood of the BA fraction exceeding 0.5%.

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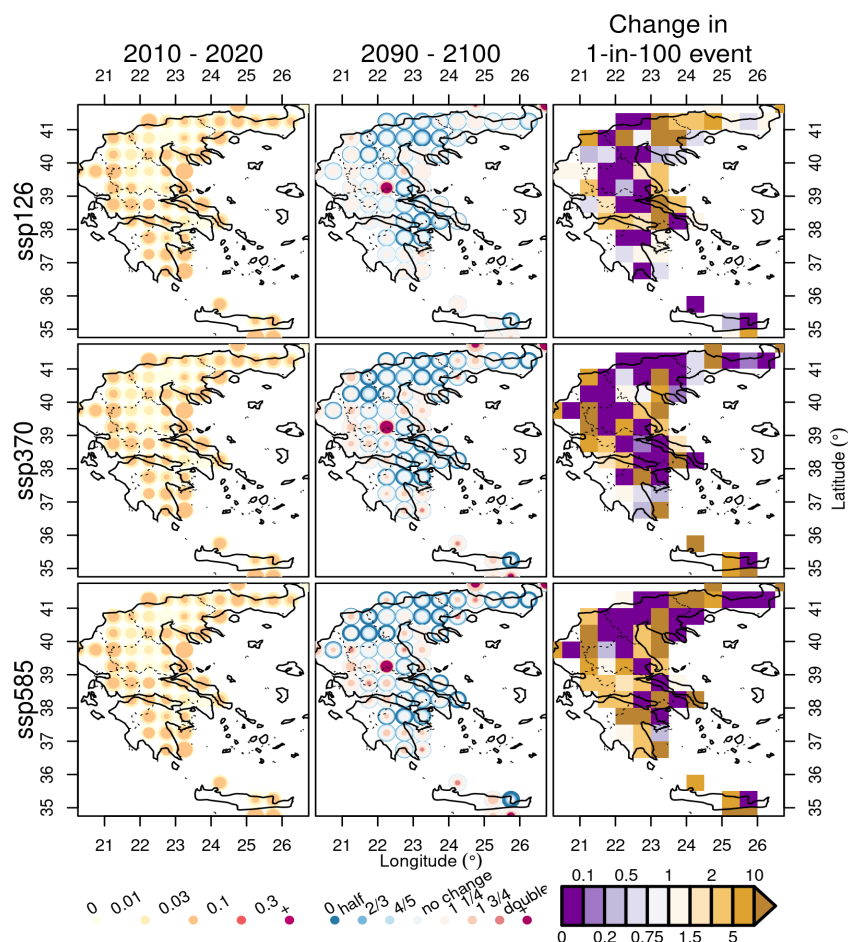
5.2.2.2 Greece

The probability of Greece experiencing BA extent similar to August 2023 (specifically, for cells with BA fraction in the 90th percentile in that month) is estimated to be 1.3% in any given year under the climate conditions of 2010-2020, according to estimates made using reanalysis data (**Table 6; Figure 17**). Bias-correction did not fully resolve all discrepancies between the GCMs and reanalysis data, and the GCMs gave a range of likelihoods spanning 0.7-1.8% for any given year under the climate conditions of 2010-2020.

In SSP126, no significant increase in likelihood of an event like 2023 is projected for any decade through 2100 (i.e., beyond the range of 0.7-1.8% for the 2010s). This is despite the likelihood of a 2023-like event in some decades being as high as 2.3 times more likely than in the 2010s under SSP126 (**Figure 17**). The lack of significance in these changes may in part reflect our strict definition of significance (i.e., no overlap with the range of the 2010s). When likelihoods vary considerably across models due to the incomplete resolution of biases, the thresholds for significance are high. The small number of observations available for model training in Greece due to its small domain size, which likely contributes to wider uncertainty bounds and higher biases than in the other, larger focal regions. Overall, Greece could see large increases in the occurrence of extreme BAs even under strong mitigation (SSP126), but uncertainties are large.

SSP350 and SSP585 show significant increases in the likelihood of an event like 2023 by the 2070s (relative to the 2010s), and also diverge significantly beyond SSP126 in the 2080s. SSP585 and SSP370 do not diverge from one another throughout this century and, in 2100, both scenarios give a likelihood of an event like 2023 of 2.5-3.3%. This range is equivalent to a 1.8-to-4.3-fold increase on the values of the 2010s, and an average return time of 31-39 years. The divergence of likelihoods between SSP126 and the two other scenarios (SSP350 and SSP585) is associated with increases in both fuel load and fuel dryness, with particularly striking differences in the latter across the scenarios (**Figure 17**).

There are some spatial patterns in the future trajectory of summer BA and the likelihood of future 1-in-100 year BA events in Greece by the 2090s (**Figure 19**). Interior parts of Greece tend to see a decline in BA, while coastal parts see an increase in BA and this pattern broadly holds for 1-in-100 year events. Across scenarios with increasingly low levels of climate change mitigation, there is an expansion to the portion of Greece's total area that experiences increased summer BA and increased frequency of 1-in-100 year events.



2112 **Figure 19:** Same as **Figure 18**, but for Greece and months July-September.

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2116 **5.2.2.3 Western Amazonia**

2117
2118 The probability of Western Amazonia experiencing BA extent similar to September-October
2119 2023 (specifically, for cells with BA fraction in the 95th percentile in that month) is estimated
2120 to be 16.58% in any given year under the climate conditions of 2010-2020, according to
2121 estimates made using reanalysis data (**Table 6**; **Figure 17**). Bias-correction did not fully
2122 resolve all discrepancies between the GCMs and reanalysis data, and the GCMs gave a range
2123 of likelihoods spanning 15.1-16.6% for any given year under the climate conditions of 2010-
2124 2020. As expected, these results suggest that the events observed in Amazonia in 2023 were
2125 not as extreme as those in Canada and Greece, with returns times of around 6-7 years.

2126
2127 We note that spatial patterns of fire are highly dependent on patterns of land use and human
2128 ignition sources in Amazonia (e.g. **Figure S14**), and an important caveat is that the BA
2129 responses to future climate and land cover change, presented below, is likely to be highly
2130 modulated by socioeconomic factors (Lapola et al., 2023; Kelley et al., 2021; Silveira et al.,
2131 2020).

2132
2133 Under the SSP585 scenario, the likelihood of an event like 2023 increases significantly in the
2134 2090s to 18.2-21.0%, representing a factor 1.2-1.3 increase versus the 2010s. This increase
2135 is primarily attributed to lower fuel moisture, in line with declines in fire weather in this region
2136 (**Section 4.2.1**). On the other hand, in the SSP126 scenario representing strong climate

2137 change mitigation, the likelihood of an event like 2023 does not change significantly at any
2138 point this century (**Table 6**).

2139
2140 The likelihood of an event like in 2023 increases substantially in the centre of the region by
2141 the 2030s under SSP370 and SSP585 (**Figure S21**) and the affected area expands through
2142 2100 (**Figure S22**). Southern regions with greater population and infrastructure densities see
2143 lesser increases in likelihood. Perhaps more important for the region's fire-sensitive forests is
2144 the projected increase in BA in SSP370 and SSP585 across forests in the north of the region,
2145 which would be expected to impact on some of Earth's most remote and pristine forests by
2146 the end of this century (**Figure S23**).

2147
2148 The 2023 event in Amazonia had a relatively high return interval compared with other focal
2149 regions. Hence we tested for change in likelihood of a rarer, 1-in-100 year event in western
2150 Amazonia. Under SSP585, we found that the likelihood of a 1-in-100 year event increased
2151 significantly by the end of the century, to 2.3-3.3% in the 2090s (representing a factor 2.2-2.9
2152 increase in likelihood). The likelihood of a 1-in-100 year event also increased significantly
2153 under SSP370 towards the end of this century, but under SSP126 no significant change
2154 occurred this century.

2155
2156 Our results show that SSPs associated with lesser climate change mitigation (SSP585 and
2157 SSP370) promote more frequent extreme BA in western Amazonia, irrespective of change in
2158 land use and human ignition factors with potential to strongly modulate the BA response to
2159 climate. There is greater potential for compounding effects of human factors and climate-
2160 driven increases in extreme BA likelihood under scenarios with no or low climate change
2161 mitigation than in the case of scenarios with high climate mitigation (such as SSP126).

2162

2163 **6 Conclusions**

2164

2165 **6.1 Summary of the State of Wildfires in 2023-24**

2166

2167 **6.1.1 Extreme Wildfire Events of 2023-24**

2168

2169 ● **Global:** A total of 3.9 million km² burned globally during the 2023-24 fire season,
2170 slightly below the average of previous seasons (4.0 million km²) and ranking 13th since
2171 2002. Despite the lower BA, fire C emissions were 16% above average, totaling 2.4
2172 Pg C, ranking 7th highest since 2003. Global C emissions were pushed up by record
2173 emissions in Canadian boreal forests and pulled down by below-average fire activity
2174 in the African savannahs (significant because fires in the African savannahs make, on
2175 average, the largest contribution of any continental biome to global mean annual BA
2176 and emissions). If global savannah fire emissions had been in line with their average
2177 in 2023-24, global fire C emissions would have been the highest on record.

2178

2179 ● **North America:** Record fire activity in Canada's boreal forests, with BA reaching six
2180 times the average and fire C emissions over nine times the average and contributing
2181 significantly towards global C emissions (24%, up from a mean value of 3% from prior
2182 fire seasons). Canada experienced extreme and widespread fires, with over 150,000
2183 km² burned, prompting evacuations of 232,000 people. Eight firefighters lost their lives.
2184 Fires in Canada led to a significant degradation of air quality, with populations living in
2185 the US and Canada exposed to atmospheric concentrations of PM_{2.5} over daily
2186 standards for periods of 2 weeks to multiple months. The United States saw generally
2187 below-average fire activity, but the Lahaina wildfire in Maui, Hawaii, resulted in 100
2188 civilian deaths, destroyed 2,000 homes, and displaced 10,000 people. Texas recorded
2189 its largest ever single fire, which destroyed 130 homes and killed two civilians.

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- **South America:** South America experienced somewhat below-average fire extent overall, but notable exceptions included significant anomalies in the northwest of the continent. In Brazil's Amazonas state, fire counts reached record highs amidst drought conditions, severely impacting air quality in Manaus. In neighbouring parts of Bolivia, Peru, Venezuela, and Guyana, also affected by regional drought, record or high-ranking levels of fire counts, extent, and carbon emissions were observed. In Chile, the Valparaíso wildfire in February 2024 killed 131 people and caused widespread property destruction.
- **Europe:** Europe experienced low fire extent in 2023-24, however the Evros fire in Greece set a new EU record for individual fire size (around 900 km²) and killed 19 people. Individual fires in Greece, Spain, Italy, Portugal, France, and Scotland led to a range of impacts including large-scale evacuations, significant suppression costs, disruption of water supplies, damage to infrastructure or agricultural lands, impacts on tourism and local economies, destruction of properties (see **Appendix A**).
- **Oceania:** Above-average fire activity and extent were observed in the savannahs, grasslands, and shrublands of western and northern Australia. Although less impactful than the 2019-20 Black Summer fires due to their remoteness, wildfires near Perth resulted in property destruction. The Tara and Mount Isa fires in Queensland destroyed 65 homes and claimed two lives. In Victoria, fires destroyed over 40 homes and injured five firefighters. In New South Wales, forest fires caused widespread smoke-related damages.
- **Asia:** Most parts of Asia saw low fire activity. Lao PDR, Thailand, and Vietnam experienced high fire counts amidst reported heatwave conditions and a possible uptick in agricultural fire use, leading to regional haze and air quality issues. In Mongolia's Dornod Aimag province, wildfires burned large parts of the Daurian steppe and firefighting was required to stop fires crossing the border into Russia.
- **Africa:** Africa experienced low fire extent in general during 2023-24, with BA 13% below average in the African grassland, savannah, and shrubland biome. However, extreme fires in Northern Africa, particularly Algeria and Tunisia, prompted significant emergency responses including assistance from the EU. In Algeria, wildfires occurring in temperatures around 50°C resulted in 34 deaths and over 1,500 evacuations, with over 8,000 personnel deployed to combat the fires. Tunisia faced similar challenges with strong winds exacerbating wildfires, leading to evacuations in the northwestern region. In coastal South Africa, fires in the Western Cape caused structure damage and evacuations.

6.1.2 Diagnosing Drivers and Assessing Predictability

- **In Canada, fire weather conditions were the primary drivers of the record-breaking fire activity and extent during 2023.** However, there were notable contributions from other drivers such as upticks in ignitions or human factors that were not explicitly represented in our analyses. This highlights potential inadequacies in predicting some ignition sources or accurately representing fire propagation in current forecasting systems.
- **The exceptional nature of Evros fire in Alexandroupolis, Greece, could not have been predicted using fire weather forecasts.** While there were discernible indications in the Fire Weather Index (FWI) records that the days around the event were more extreme than most days in the fire season, similar conditions were also observed in other periods without resulting in the same catastrophic impacts. This

2245 highlights the intrinsic difficulties in forecasting isolated extreme fire events and
2246 underscores the need to advance early warning systems beyond fire weather to
2247 consider fuel availability and ignition variability.

- 2248
- 2249 ● **The extreme fire season in western Amazonia was driven by prolonged drought**
- 2250 **conditions linked to the strong El Niño.** Many fires developed, triggered by lightning
- 2251 ignitions early in the season amidst high fire weather anomalies. Other than weather
- 2252 conditions that acted as persistent controls for fire activity, several periods with high
- 2253 active fire counts were not predicted in late August and throughout September likely
- 2254 due to the result of unrepresented ignition sources.
- 2255
- 2256 ● **In all focal events, extreme burned areas were driven by anomalies in multiple**
- 2257 **controls on fire. Weather, fuel moisture, and fuel abundance were all critical**
- 2258 **factors.** The synchronous occurrence of anomalies in fire weather, high fuel load, and
- 2259 low fuel moisture created the conditions leading to the anomalous burned areas
- 2260 recorded in all three events. This underscores that no single bioclimatic factor can
- 2261 explain the most severe fires. Instead, multiple contributing controls must coincide for
- 2262 the most extreme events to arise. Factors such as ignitions, suppression, and landscape
- 2263 fragmentation, related to human activities, likely played important roles in modulating the
- 2264 Western Amazonian and Greece events.
- 2265
- 2266 ● **Fuel load is an important modulator of the relationship between fire extent and**
- 2267 **fire weather.** Higher and/or drier fuel loads combined with high fire weather conditions
- 2268 caused the unprecedented extent of burning in Canada and Western Amazonia. The
- 2269 boundaries (extinction points) of extreme fires in Canada and Greece often
- 2270 corresponded to areas with lower fuel loads, demonstrating that discontinuity in fuel
- 2271 availability constrained fire spread.
- 2272

2273 6.1.3 Attribution to Global Change

- 2274
- 2275 ● **In Canada, anthropogenic forcing increased the chance of high fire weather in**
- 2276 **2023. Total climate forcing led to higher BA, whereas socio-economic factors**
- 2277 **may have decreased burning.** Anthropogenic forcing (resulting from greenhouse gas
- 2278 emissions and land-use change) at least doubled the probability of experiencing high
- 2279 fire weather in June 2023. It is virtually certain that total climate forcing (resulting from
- 2280 climate change since the pre-industrial period) increased the BA in Canada by up to
- 2281 40.1%. Socioeconomic factors related to land use, ignitions and suppression may have
- 2282 reduced burning, though with low confidence.
- 2283
- 2284 ● **In Greece, anthropogenic forcing increased the chance of high fire weather in**
- 2285 **2023. Total climate forcing led to higher BA, and socio-economic factors may**
- 2286 **have increased or decreased BA.** Anthropogenic forcing increased the probability of
- 2287 experiencing high fire weather in August 2023 by 1.9-4.1 times. It is likely that total
- 2288 climate forcing increased the BA in Greece by up to 17.7%, whereas socio-economic
- 2289 factors could have led to an increase or decrease. Climate change has increased
- 2290 average BA in the wider Mediterranean region, but this has been mainly offset by
- 2291 socio-economic factors.
- 2292
- 2293 ● **In western Amazonia, anthropogenic forcing has greatly increased the chance**
- 2294 **of high fire weather in 2023. It is virtually certain that total climate forcing led to**
- 2295 **higher BA, and very likely that socio-economic factors also contributed.**
- 2296 Anthropogenic forcing increased the probability of experiencing high fire weather by
- 2297 more than a factor of 20. It is virtually certain that total climate forcing increased BA in
- 2298 western Amazonia by up to 49.7%, and very likely that socio-economic factors

2299 exacerbated the increase. Climate change has increased today's average BA in the
2300 region, and all forcings have led to an overall increase in burning.
2301

2302 6.1.4 Seasonal and Multi-Decadal Outlook

- 2303
- 2304 ● **While the positive El Niño phase of ENSO is subsiding towards a neutral phase**
2305 **in 2024, the Indian Ocean Dipole is persisting in its positive phase, with potential**
2306 **to influence global fire patterns.** The 2023–2024 El Niño event emerged as the
2307 fourth-largest positive anomaly on record, however most simulations forecast a
2308 transition to ENSO-neutral conditions in 2024. Positive IOD phases are associated with
2309 elevated BA Amazonia, Indonesia, and parts of Australia though these tend to depend
2310 on interactions with ENSO.
2311
- 2312 ● **Seasonal predictions of fire weather through August 2024 highlight moderate**
2313 **positive anomalies (75th percentile) in parts of Canada and much of South**
2314 **America including Amazonia,** as well as in central Africa, Southern Africa, southeast
2315 Europe, western Australia, southeast Asia, and northeast Asia. Extreme (95th
2316 percentile) anomalies are rare in the forecast through August 2024.
2317
- 2318 ● **In Canada, future likelihood of events like in 2023 are increased by rising fuel**
2319 **load and dryness, with high mitigation pathways significantly reducing these**
2320 **risks.** The likelihood of extreme fire events similar to those in June 2023 is currently
2321 low (0.15% for any given year in the 2010s, or 1-in-700 years). The likelihood is
2322 expected to increase to 0.42-2.2% by the 2040s, and thereafter continue to increase
2323 under high emissions scenarios (SSP585) reaching 2.1-3.7% by the 2090s. The
2324 probability of experiencing such events at least once by the 2090s is estimated to be
2325 59-87% under SSP585, compared to 18-73% under SSP126. Individuals born in
2326 Canada during the current decade have a 54-87% likelihood of witnessing an event on
2327 the scale of 2023 again in their lifetime under SSP370 (a scenario close to present day
2328 trajectories without additional mitigation efforts).
2329
- 2330 ● **In Greece, high mitigation scenarios result in no significant change in the**
2331 **likelihood of an event like 2023 through 2100, whereas scenarios without**
2332 **mitigation lead to significant increases.** The current likelihood of extreme fire events
2333 like those in August 2023 is around 1.3% annually for any given year in the 2010s (or
2334 roughly 1-in-100 years). Under SSP126, there is no significant increase projected
2335 through 2100 whereas significant increases in likelihood (to up to 3.3%) are projected
2336 under SSP585 and SSP370 by the 2090s, implying more frequent extreme fire events.
2337 Coastal areas of Greece are expected to experience the greatest increases in risk,
2338 whereas interior regions may experience declines.
2339
- 2340 ● **In Western Amazonia, the effects of climate change on extreme fire likelihood in**
2341 **future scenarios is chiefly governed by fuel moisture effects, which are likely to**
2342 **be strongly modulated by local changes in land use and human ignition sources.**
2343 **The potential for compounding effects between fuel dryness and local stress**
2344 **factors is minimised in high-mitigation scenarios.** The current likelihood of extreme
2345 fire events like those in September-October 2023 is 16.6% (or roughly 1-in-6 years).
2346 Under SSP585, the likelihood of such events increases significantly to 18.2-21.0% by
2347 the 2090s. In addition, there is a significant rise in the probability of 1-in-100 year
2348 events by the end of the century under SSP585 and SSP370. No significant rise in
2349 probability of 1-in-6 or 1-in-100 events is seen under SSP126. Hence, while increased
2350 fire risks related to climate change in the Amazon can be compounded by human
2351 activities, this is least likely under SSP126. The impact of extreme fire events is
2352 expected to be severe in pristine northern forests, emphasising the need for strong
2353 climate change mitigation.

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- Overall, high-emissions scenarios (e.g. SSP585, SSP370) lead to significantly increased likelihood of major events like those seen in the 2023-24 fire seasons in future, highlighting the critical importance of strong climate change mitigation efforts (e.g. SSP126) to reduce the future likelihood of extreme fire events. Even though SSP370 and SSP585 scenarios suggest significant increases in extreme events, the confidence that strong mitigation (SSP126) can avoid significant portions of increased risk is a major promising outcome of this work. Our findings emphasise the importance of continued and enhanced mitigation efforts.

6.2 Roadmap for the State of Wildfires Report

The report successfully achieved its primary objectives, which included identifying and contextualising extreme wildfires and wildfire seasons over the past year, selecting focal events with significant societal and environmental impacts, and diagnosing the factors contributing to these events. The report also assessed the predictive capabilities of existing systems, highlighting their strengths and limitations, and attributed the occurrence of focal events to anthropogenic factors, including climate change and land use. Additionally, it provided an outlook for future wildfire probabilities, emphasising the limitations of current long-term forecasting tools and the increasing likelihood of extreme fire events under future climate scenarios, particularly highlighting the need for improved accuracy in regional projections.

The fire science community is currently navigating several research frontiers to improve prediction of extreme fires and understanding of their causes, with the view to enhance preparedness, response, mitigation, and adaptation to wildfires in wider society. The field is advancing its ability to observe individual fires, assess conditions leading to extreme fires, and predict their occurrence on timescales ranging from hours to decades. Additionally, there is increasing focus on monitoring and modelling the diverse impacts of extreme fires on society, the environment, and the economy.

As part of this inaugural edition of the State of Wildfires report, we present, in **Appendix B**, a stocktake of current capabilities, challenges, and emerging opportunities in the observation and modelling of extreme fires and their impacts. **Appendix B** is intended for the interdisciplinary community of fire scientists and represents a contribution to agenda-setting within this field of research. It will not be revised annually but may be revisited in future to serve as a stocktake of progress in this field. Here, we briefly summarise the specific role that the State of Wildfires report should serve as advances in the observation, prediction, and modelling of extreme fires and their impacts come to fruition.

6.2.1 Definition

The State of Wildfires Report will facilitate a community effort on a protocol for defining extreme fire events or fire seasons.

This report emphasises important issues in the definition of extreme fires, a problem that affects the definition of many terms used in fire science including ‘wildfire’ and ‘megafire’ (**Appendix B**; e.g. Shuman et al., 2022, Linley et al., 2022). Definition is complicated by the impact of fires on society and the environment across many impact sectors, with the magnitude of impact not necessarily correlating with observable fire traits such as BA (**Appendix B**). In future years, regional experts would benefit from a protocol or guidelines that can be used for categorising extreme fire events or seasons. To support future iterations of the State of Wildfires report, we will coordinate workshops with broader sections of the fire science community with the aim to produce guidance for future years. Central to this task is

2407 the inclusion of communities from broad geographies so that any output respects fire impacts
2408 that are considered to be regionally significant.

2409

2410 **6.2.2 Observation**

2411

2412 **The State of Wildfires report will advocate for and utilise new harmonised fire**
2413 **observation products.**

2414

2415 Consistent, long-term records of fire extent and properties are fundamental for studying
2416 extremes, which cannot be characterised without reference to historical ranges (**Appendix B**).
2417 The MODIS instrument has been crucial in tracking global fire progression and emissions over
2418 two decades, but its continuity is threatened as the Terra satellite nears decommissioning,
2419 necessitating harmonisation with newer datasets like VIIRS, Landsat, and Sentinel with
2420 MODIS records for consistent fire observation.

2421

2422 The State of Wildfires report further underscores the critical strategic need for a continuous
2423 and harmonised dataset of fire observations beyond the MODIS era (**Appendix B**). To support
2424 future iterations of the report, we will advocate for the provision of harmonised products within
2425 the Earth observations communities. In addition, regional products often provide scope to
2426 characterise the extremity of events over multi-decadal timescales and are now being provided
2427 in globally harmonised formats compatible with global analyses such as ours. These regional
2428 dataset should be utilised in future iterations of the State of Wildfires report.

2429

2430 **The State of Wildfires Report will stimulate progress on combining multiple fire**
2431 **observation streams to better identify and characterise extreme fire.**

2432

2433 This report highlights the need to advance our capacity to observe fires that are impactful in
2434 diverse ways (**Appendix B**). In particular, there is a growing need to move ‘beyond burned
2435 area’ and towards a wider set of intensity, severity and behaviour metrics that often relate
2436 more strongly to impacts on society and the environment than BA. The integration of individual
2437 fire data from the Global Fire Atlas in this iteration of the State of Wildfires Report is one
2438 example of including wider fire parameters such as size and rate of growth. Further
2439 applications of the dataset or other individual fire atlases (e.g. Laurent et al., 2018; Artés et
2440 al., 2019) could include assessing days with many synchronous large fires, which challenge
2441 fire management (e.g., Abatzoglou et al., 2021), and identifying impactful fires by their
2442 intersection with population centres (e.g., Modaresi Rad et al., 2023). Combining individual
2443 fire behaviour data with fire radiative power and biomass combustion estimates might better
2444 identify intense or severe fires with significant consequences for ecosystems and society (e.g.,
2445 Nolan et al., 2021a).

2446

2447 Overall, the State of Wildfires report must stimulate progress on moving ‘beyond burned area’
2448 and combining diverse observational capacities to better identify and characterise extreme fire
2449 events, and we intend to expand our use of such insights in future iterations. Likewise, the
2450 report must be poised to adopt into its definition of extreme fires any emerging datasets that
2451 quantify fire impacts on the various impact sectors outlined in **Appendix B**.

2452

2453 **6.2.3 Prediction**

2454

2455 **The State of Wildfires report will advocate for the use of extended range forecast to**
2456 **identify early onset of fire weather conditions.**

2457 Global fire danger monitoring systems currently use short to medium-range weather forecasts,
2458 typically up to 10 days (**Appendix B**). However, state-of-the-art seasonal forecasting systems
2459 can predict fire-conducive conditions up to one month in advance, and in some regions, up to

2460 two months. Longer-term predictability is achievable in regions where fire activity corresponds
2461 strongly with climate modes such as ENSO.

2462 By presenting an annual opportunity to take stock of current capabilities in forecasting horizons
2463 for significant global fire events, the State of Wildfires report will continue to showcase and
2464 advocate for advances in the forecasting window for extreme fire potential on subseasonal to
2465 seasonal timescales.

2466 **The State of Wildfires report will stimulate progress on the use of AI and informatics**
2467 **methods to aid the forecast of fire activity globally.**

2468 There is strong potential for data-driven applications, such as machine learning, to improve
2469 predictions of extreme fire occurrence and move beyond traditional prediction systems based
2470 on meteorological indices such as FWI (**Appendix B**). These methods can incorporate diverse
2471 data inputs representing the influence of fuel loads, fuel moisture, ignition opportunities, and
2472 suppression on fire likelihood, therefore improving upon indices that are mostly a function of
2473 weather conditions. Tools used in this report, such as ConFire, are structured to harness new
2474 data as they become available, including near-real-time data, improved fuel observations, and
2475 detailed human/fire interaction data. ConFire is also being developed to optimise its
2476 representation of extreme fire events by incorporating more flexible response curves into its
2477 BA predictions (Barbosa 2024). The State of Wildfires report will showcase the benefits of
2478 these emerging technologies in enhancing fire prediction and management.

2479 **6.2.4 Attribution**

2480
2481 **The State of Wildfires report will promote the enhancement of low-latency attribution**
2482 **approaches.**

2483 Fire attribution techniques are relatively novel compared with more established approaches
2484 for extremes such as heatwaves. Part of the challenge in attributing fire is that it is a complex
2485 hazard comprising multiple compound risks across both meteorological and human drivers,
2486 all of which must be represented in the driving data sets used by models (**Appendix B**). A
2487 particular challenge that we faced in our current work was latency in the reanalysis datasets
2488 used to drive our model for novel attribution of fire extent. Our working group will therefore
2489 engage with the ISIMIP project to promote the creation of low-latency reanalysis products to
2490 support responsive attribution assessment in future State of Wildfires reports, and in other
2491 near real-time applications. An additional avenue for enhancement of future reports is to
2492 include a greater number of climate models in the attribution work by widening the participation
2493 of other groups working on attribution internationally.

2494 **The State of Wildfires report will support the expansion of attribution methods that**
2495 **target a range of extreme fire metrics.**

2496 Extending attribution approaches to a broader range of extreme fire properties (e.g. aspects
2497 of the fire size or fire behaviour distributions) is increasingly possible as observations of
2498 individual fire characteristics and behaviour avail (**Appendix B**). Capacity to attribute
2499 individual fire properties can be built by coupling the ConFire model with process-based
2500 models within attribution frameworks, allowing attribution of specific extreme fire
2501 characteristics to climate change and other forms of global change.

2502 2503 **6.2.5 Projection**

2504
2505 **The State of Wildfires report will harness projections from multi-model ensembles.**

2506 In the coming years, FireMIP and ISIMIP will provide ensembles of model projections of future
2507 BA for the first time (**Appendix B**). The State of Wildfires report will make use of these
2508 simulations as soon as they are available, thus improving upon the single model (ConFire)
2509 employed in the current edition and improving characterisation of uncertainty in the
2510 projections.

2511 **The State of Wildfires report will support the expansion of projection methods to target**
2512 **a range of extreme fire metrics.**

2513 An ambition for the State of Wildfires report is to provide projections of future fire properties
2514 “beyond burned area”, such as the fire size and behaviour distribution (**Appendix B**). This
2515 capacity can be built by combining the ConFire model with multiple process-based models
2516 from FireMIP and ISIMIP. Advances in this area have particular societal relevance in resource
2517 planning, as changes in fire behaviour (not only fire extent) are likely to factor into decision-
2518 making around future suppression resource requirements.

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7 Appendix A: Year in Review by Continent

This appendix includes the review completed by an expert panel to supplement our quantitative analyses of extremes in the 2023-24 fire season (see **section 2.1**). Details of the assembled panel are provided in **Table A1**, below.

Table A1: Experts contributing to the identification of extreme events and characterisation of the global fire season during March 2023-February 2024.

Region	Expert	Country of Organisation / Nationality	Professional Background(s)	Others Consulted
Africa	Natasha Ribeiro	Mozambique	Research	Robert Ang'ila, Karatina University; Kebonyethata Dintwe, Botswana University; John Mendelsohn, Okavango Research Institute; Ronald Heath, Forestry South Africa; Helen De Klerk, Stellenbosch University
	Sally Archibald	South Africa	Research	
Asia	Bambang Saharjo	Indonesia	Research, Litigation	
	Veerachai Tanpipat	Thailand	Research, Fire Control and Management Instructor and Consultant	
Europe	Paulo Fernandes	Portugal	Research	Davide Ascoli, University of Turin, Italy; Hellenic Agricultural Organization "DIMITRA"; Institute of Mediterranean Forest Ecosystems; Niall MacLennan, Scottish Fire and Rescue Service
	Stefan Doerr	UK / Germany	Research	
	Gavriil Xanthopoulos	Greece	Research	
North America	Crystal Kolden	USA	Research, Firefighting	
	Jacquelyn Shuman	USA	Research	
	Piyush Jain	Canada	Research	
Oceania	Hamish Clarke	Australia	Research, Environmental Management	Grant Pearce, Fire and Emergency New Zealand; Simeon Telfer, South Africa Country Fire Service; Agnes Kristina, Department of Fire and Emergency Services; Russell Stephens Peacock, Queensland Fire and Emergency Services; Chris Collins, Tasmania Fire Service; David Field, NSW Rural Fire Service.
	Sarah Harris	Australia	Research, Emergency Management	
South America	Dolors Armenteras	Colombia	Research	The Chico Mendes Institute for Biodiversity Conservation (ICMBio), Brazil, Santarém Office
	Liana Anderson	Brazil	Research, Disaster risk reduction strategies	

2527
2528

7.1 Africa

South Africa and Botswana experienced higher than average BA and fire size (**Figure 2, Figure 4**) in 2023-24, which some regional experts had expected following three consecutive years of above-average rainfall that increased grassy fuel loads in the fuel-limited savannas and grasslands. This has potentially been exacerbated by a lack of prescribed burning and active fire suppression in the privately held land and conservancies in the region, that likewise would have resulted in fuel build-up (*Atlas of Namibia*, 2021). The socio-economic impacts of these large fires were minimal (extensive grassland fires linked to high rain years are expected due to periodic 7-20 year wet-dry cycles in these ecosystems).

In East Africa, the area burned was extremely low in 2023-24. This was in line with the expectations of regional experts given the effects of a triple La Niña in this region, which causes droughts in East Africa (in contrast to southern Africa). This multi-year drought meant that there were limited grass fuels to burn and reduced the likelihood of spread of accidental ignitions in many of the East African rangelands. However, the extreme fire weather enabled fires to burn through upland forests, which are not normally flammable. This included a regionally significant fire in Aberdare Forest, Nyeri county, Kenya, which reportedly burned 160 km² on 17th of Feb 2023 (*Citizen Digital*, 2023).

The 2023 heatwave in North Africa exacerbated fire behaviour in the region (*Al Jazeera*, 2023a). Algeria recorded significant fires in the latter half of July, facilitated by high temperatures that reached upwards of 48°C (*Al Jazeera*, 2023b). Over 8,000 personnel, including firefighters and the military, were deployed to combat rapidly spreading fires across 15 provinces (*South African Broadcasting Corporation*, 2023a). These efforts were critical in managing fires that forced over 1,500 people from their homes (*euronews*, 2023). Despite these efforts, the wildfires claimed the lives of at least 34 individuals, including 10 soldiers (*Al Jazeera*, 2023b).

Neighbouring Tunisia also faced wildfire outbreaks, exacerbated by strong winds that carried fires across the national border from Algeria, leading to the closure of two border crossings (*Reuters*, 2023a). The Tunisian wildfires prompted evacuations in the north-western region of Tabarka, affecting at least 300 people and extending firefighting efforts to Bizerte, Siliana, and Beja (Sullivan and Tondo, *The Guardian*, 2023). Resources such as firefighting aircraft and personnel were sent from EU nations to help tackle the fires, despite the challenging conditions imposed by near-record temperatures of 49°C (Gauldie, *AirMed&Rescue*, 2024). In August, forest fires in mountainous regions of Morocco were also fanned by strong winds and facilitated by protracted hot spring and summer temperatures (Erraji, *Morocco World News*, 2023; Copernicus Climate Change Service, 2024a).

During December 2023-January 2024, the Western Cape of South Africa experienced wildfires related to prolonged hot and windy conditions, causing substantial damage and prompting widespread evacuations. In the Overstrand local municipality, which includes coastal towns like Pringle Bay and Betty's Bay, multiple fires necessitated evacuations and destroyed properties. The Hangklip area between Pringle Bay and Betty's Bay was particularly affected with the fires destroying properties in the Sea Farm private nature reserve. On January 29, a "code red" status was declared, indicating a serious threat to residential areas, and evacuations were advised for communities including Silversands and Seafarms (*Crisis24*, 2024; *AfricaNews*, 2024). A wildfire swept from Simonstown to Scarborough in Cape Town, necessitating large-scale evacuation (Hough, *IOL News*, 2023). This fire was challenging due to its rapid spread fueled by strong south-easterly winds and high temperatures. The firefighting efforts were supported by multiple helicopters and ground teams (*South African Broadcasting Corporation*, 2023a). The most extensive damage was reported from the Kluitjieskraal fire near Wolseley, where over 220 km² were burned, and more than 40 structures were destroyed. This fire also prompted evacuations and remained uncontained for

2584 several days due to its size and complex terrain that hindered ground access (*Crisis24*, 2024)
2585 . Despite these extreme wildfires, the plantation forestry industry was not affected, with
2586 relatively low losses due to fire.
2587

2588 7.2 Asia

2589
2590 The 2023-24 fire season in Asia was generally not an extreme one, with much of central Asia
2591 experiencing low BA. Siberia, which has seen several record-breaking fire seasons since 2020
2592 resulting in globally significant fire emissions (Zheng et al., 2021), also experienced a
2593 somewhat typical year for BA and fire C emissions. Likewise most provinces of China and
2594 states of India experience a fairly typical fire season.
2595

2596 Nonetheless, there were regional examples of high fire activity in the 2023-24 fire season. The
2597 Dornod Aimag province of eastern Mongolia, near the borders with Russia and China,
2598 experienced several extreme fires during April 2023 that are also visible as anomalies in the
2599 global fire observations (**Figure 2, Figure 4**). Over 15% of the area of Dornod Aimag burned
2600 in 2023-24 in contrast to the 23-year average of below 5%. The province includes the
2601 Mongolian part of the Daurian steppe, notable for being one of the last remaining undisturbed
2602 steppes in the world (UNESCO World Heritage Centre, 2017). Unusually dry and warm
2603 conditions in eastern Mongolia during spring led to severe wildfires. A notable wildfire spread
2604 into Dornod from neighbouring Sukhbaatar province, fanned by windy, dry conditions (*Borneo
2605 Bulletin*, 2023). The National Emergency Management Agency mobilised over 250 individuals,
2606 including firefighters and local residents, and helicopters were deployed to manage the fast-
2607 spreading fires (*Borneo Bulletin*, 2023). The effects of these wildfires on herder and nomadic
2608 populations and the Mongolian Red Cross has provided aid to 4,800 herder households
2609 (International Federation of Red Cross and Red Crescent Societies, 2023).
2610

2611 Although BA extent and fire counts were overall below the 2002-2023 average along the
2612 southern border regions of Russia during 2023-24 (**Figure 2, Figure 4**), a number of disruptive
2613 wildfires fanned by strong winds broke out during April and May and affected regions bordering
2614 Kazakhstan, such as in the Tyumenskaya, Omskaya, and Amurskaya Oblasts, and Mongolia,
2615 such as in Irkutsk and Krasny Yar where at least one fatality was recorded (*Le Monde*, 2023).
2616 As well as detecting anomalies in fire size and rate of spread in these areas, the Global Fire
2617 Atlas also identified regionally large and fast-moving wildfires in the Russia-China border zone
2618 of Manchuria (**Figure 4**), however these were not widely reported on by media outlets or local
2619 authorities.
2620

2621 Lao People's Democratic Republic (PDR) experienced a notable fire season in 2023-24,
2622 marked by record-setting BA at national level since 2002 in the MODIS BA data (**Figure 2;
2623 Figure S6**). The fires were widespread, affecting various provinces from the south to the north,
2624 including Attapu, Khammouan, Louangphabang, Xaignabouli, and Bokeo. In Attapeu, BA in
2625 2023-24 was over twice the average of prior fire seasons since 2002. The fires in 2023-24
2626 were generally small in scale but anomalously numerous, consistent with the widespread use
2627 of slash-and-burn agricultural fires in these regions that have been problematic for regional air
2628 quality in this region during recent years (Meadley, *Laotian Times*, 2024). The uptick in fire
2629 counts in 2023-24 has been attributed in part to economic factors such as the high price of
2630 cassava and demand for greater corn supplies to supply animal feed, which act as incentives
2631 for farmers to clear forests for additional planting (Bhandari, *Radio Free Asia*, 2024). On top
2632 of economic factors, a heatwave that spanned South and Southeast Asia in April 2023 was
2633 reported to have been an enabling driver (Zachariah et al., 2023). The persistent smoke from
2634 these fires worsened air quality significantly in southeast Asia, where efforts to manage
2635 transboundary haze have been challenging during regional droughts, despite a new
2636 transboundary agreement being signed in 2023 (Antara News, 2023). Differences in fire
2637 management between Thailand and Lao PDR were evident during the 2023 event, with

2638 authorities intensifying patrols and seeking to control forest fires and agricultural burning for
2639 improved air quality in Thailand (Meadley, *Laotian Times*, 2024). Conversely, deforestation
2640 remains a critical issue in Lao PDR, with the Laotian government facing challenges in gaining
2641 local community support for the prevention of agricultural expansion and logging.

2642
2643 Earth observations data showed high-ranking BA anomalies and fires with large size and rate
2644 of growth during 2023-24 in several regions of Pakistan, Iran, Iraq, and parts of the Levant
2645 region (**Figure 4**), consistent with reports of extreme drought-driven wildfires in some of these
2646 regions (*Reuters*, 2023b).

2647

2648 **7.3 Europe**

2649

2650 Overall, fires burned 8,400 km² in Europe from March 2023 to February 2024 according to the
2651 European Forest Fire Information System (EFFIS, 2024), of which 64.5% were from July to
2652 September and 18.1% were in March and April. Large fires (>5 km²) amounted to 53.4% of
2653 the total BA, and those particularly large (>100 km², n=5) accounted for 17.7% of the total
2654 burned area. More than half (52.6%) of the BA corresponded to transitional woodland, with
2655 forests, shrublands and grasslands, and agriculture respectively amounting to 19.1%, 13.2%,
2656 and 14.4%. At least 44 people died as a direct result of wildfires (Copernicus Climate Change
2657 Service, 2024; Centre for Research on the Epidemiology of Disasters, 2024).

2658

2659 Most countries in the Mediterranean Basin experienced mild to typical fire seasons in general,
2660 with variable timing but affecting mostly non-forest (open) vegetation types (EFFIS, 2024). In
2661 the Balkans, fire activity varied among countries, but was mostly very low by historical patterns
2662 such as in Croatia, however, a major exception was Greece described in more detail below
2663 (**Figure 2, Figure 4, Figure A1**). The other exceptions were North Macedonia, with a typical
2664 fire year and Bulgaria, the worst year in a decade, with fire activity extending into October in
2665 both countries; and Bosnia and Herzegovina, Serbia and Montenegro, where collectively ~270
2666 km² burned in January-February 2024.

2667

2668 Greece's 2023 fire season was reviewed at length by Xanthopoulos et al. (2024). It was the
2669 second worst on record regarding total area burned (1,727 km²), despite the recent efforts to
2670 strengthen the firefighting mechanism of the country with more aerial resources and new
2671 personnel, after another challenging fire season in 2021. The situation was kept under control
2672 until mid-July, but in the period July 13-27, maximum temperature in many parts of the country
2673 exceeded the average for the 2010-2019 period by as much as 10°C, according to the records
2674 of the National Observatory of Athens. This resulted in multiple fire starts pushing the limits of
2675 firefighting, which relies heavily on the aerial resources. The fires starting 18th of July on the
2676 tourist island of Rhodes grew rapidly on the second day, finally burning 207 km² and stopping
2677 at the sea. About 20,000 tourists had to be evacuated from hotels along the coast. While the
2678 fire on Rhodes was still burning, three forest fires started on 3rd of July started near the city
2679 of Aigio, in North Peloponnese, on the island of Corfu and near the town of Karystos in the
2680 south of Evia island. On July 25th a Canadair CL-215 crashed near the village of Platanistos
2681 while fighting this last fire. Then, on July 26th, the tail of a cold front that passed over Greece,
2682 with the characteristic wind direction change that accompanies it, created further challenges,
2683 as a number of fire starts in central Greece and Thessaly spread fast, burning mostly in light
2684 fuels, challenged firefighters and threatened inhabited areas. One of the fires entered an Air
2685 Force base on the 27th, causing a powerful explosion of ammunition that resulted in damages
2686 to the town of Nea Aghialos 5 km away. By the end of July, the BA across the country had
2687 reached 550 km².

2688

2689 The next wave of multiple challenging fires in Greece began on 19th August. A lightning-
2690 caused fire that started before dawn NE of Alexandroupolis in the prefecture of Evros received
2691 limited attention at first and was destined to become the largest fire in recent European history.

2692 The fact that firefighting resources were focussed on evacuation of the villages in the path of
2693 the fire, rather than fire suppression may have contributed to its eventual size. On the 21st of
2694 August, a second fire started to the north of the first one, near the village of Dadia. Fanned by
2695 a strong NE wind, it spread quickly and within a few hours it reached the rear of the first one.
2696 On that day fire behaviour both in terms of spread and intensity was extreme (Athanasiou,
2697 2024). Nineteen migrants were trapped by the flames and were found dead on the 22nd of
2698 August. Another group was saved by the firefighters at the last moment. The authorities
2699 emphasised safety and evacuated the hospital in the outskirts of Alexandroupolis.

2700
2701 On Aug. 22, while the Evros fire was the focus of attention, a fire originating at more than one
2702 point near the village of Phyli, south of mount Parnis in Attica, at the outskirts of Athens started
2703 growing against the strong NE wind. Once more, many settlements were evacuated and
2704 firefighting attention focused on protecting homes, as the fire moved slowly up the mountain
2705 slopes finally burning 62 km² in three days. The Evros fire kept growing at various rates for
2706 the next 15 days, finally reaching 938 km² and becoming the largest on record in recent history
2707 in Europe. The simultaneous spread of the Evros fire, the fire in Attica and a number of smaller
2708 fires, is likely to have increased the growth rate statistics (km day⁻¹) for fires in the region
2709 (**Figure 4** and **Figure A1**).

2710
2711 The BA of the Evros fire included 258 km² of deciduous oak forest and 218 km² of oak forest
2712 mixed with other species (Konstantinos Kaoukis, personal communication). The usually most
2713 challenging forest types regarding fire behaviour, contributed less to burned area: 128 km²
2714 forest and 152 km² of evergreen shrubs. The fire was mostly brought into control only when it
2715 reached agricultural areas and barren lands. The final size of the Evros fire may not be solely
2716 attributed to adverse meteorological conditions. One aggregating factor may have been the
2717 recent shift in directing firefighting personnel more strongly from suppression towards
2718 evacuation, another the emphasis on aerial firefighting resources (Xanthopoulos et al., 2024).
2719 The latter was not effective once the extreme fire behaviour commenced (21st to 23rd of
2720 August). Deep forest litter layers further hampered fire suppression in some areas, although
2721 a group of local forest workers working with handtools, were credited by the local forest service
2722 officers with control of a large part of the fire perimeter to the north, saving an estimated 100
2723 ha of forest (Athanasiou, 2024).

2724
2725 Italy was the second most affected country after Greece, with continuous fire activity from July
2726 to October. More than 1,000 km² burned in the country, of which 69% were in Sicily (including
2727 17 fires >10 km²), although the largest fire (31 km²) occurred in the nearby region of Calabria
2728 (Istituto Superiore per la Protezione e la Ricerca Ambientale, 2023). A defining characteristic
2729 of these large fires was the importance (42% overall) of agricultural land in the BA composition.
2730 The outskirts of Palermo and the Madonie Natural Regional Park were impacted by multiple
2731 wildfires in late September, causing one fatality and affecting wildland-urban interfaces, farms,
2732 and tourism.

2733
2734 Fire activity was insignificant in France, except for benign mountain burning (175 km²) in
2735 March-April and then in January-February, mostly in the western Pyrenees. Like in France,
2736 the north of Spain (Asturias-Cantabria) experienced unusual Spring burning activity,
2737 amounting to 423 km² during late March and early April (Educación Forestal, 2023a). In
2738 particular, the Foyedo wildfire (27th March 2023) was the largest on record for Asturias,
2739 burning 101 km² across variable vegetation but with the predominance of conifer plantations,
2740 mostly *Pinus pinaster*. It was a wind- and spot-driven fire but its soil and overstorey burn
2741 severity were respectively low and mostly moderate, as slower-drying fuels were not available
2742 to burn (Cátedra Cambio Climático de la Universidad de Oviedo, 2023).

2743
2744 Only two other notable large wildfires occurred in 2023 in continental Spain, and again were
2745 unusual in that they happened in spring rather than summer. The Villanueva de Viver wildfire
2746 (23rd March 2023, Castellón and Teruel) burned around 50 km² and was driven by abnormal

2747 seasonally dry conditions, combined with a shift in wind direction. It mostly burned naturally-
2748 regenerated continuous pine forest of *Pinus halepensis*. Canopy fire severity was
2749 heterogeneous, with 39% of the wildfire area being classified as high to very high severity
2750 (Mediterranean Center for Environmental Studies, 2023). The cost of fire suppression was
2751 2M€ and 1,800 people were evacuated (*Las Provincias*, 2023).

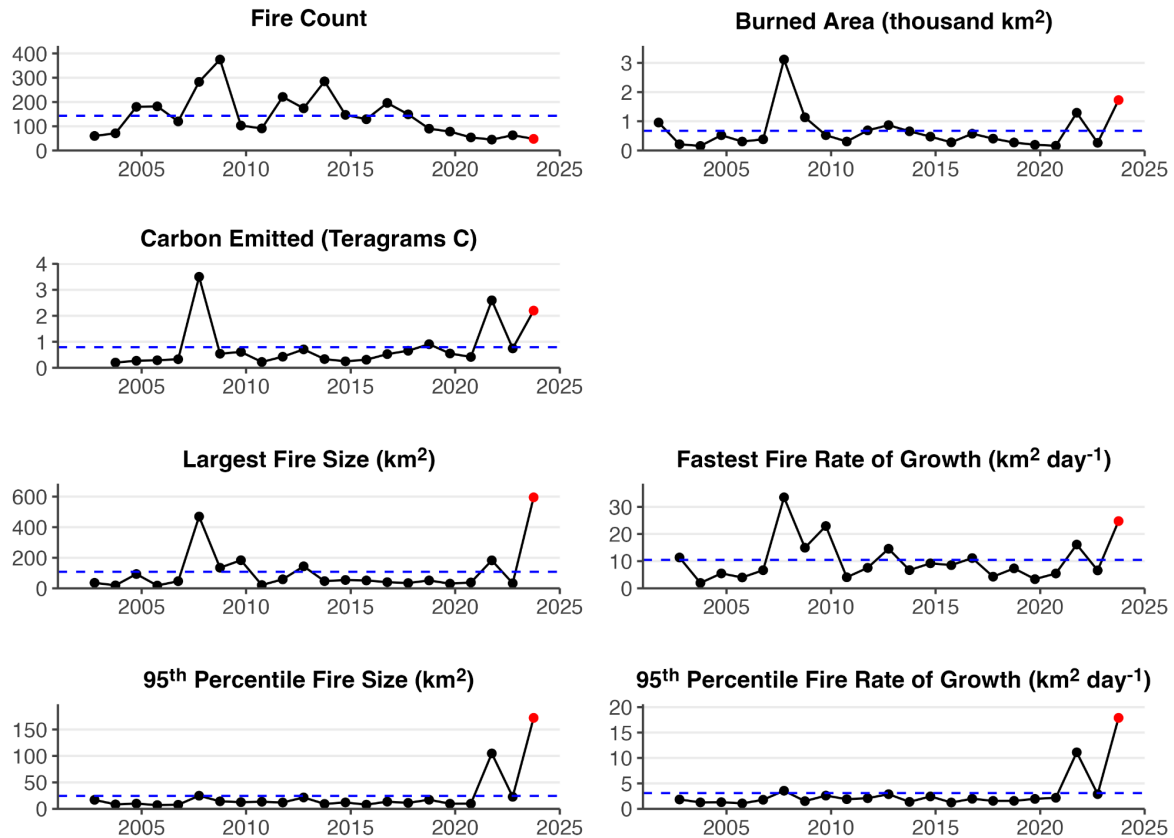
2752
2753 At over 100 km², the Pinofranqueado wildfire (17th May 2023, Cáceres) was the largest fire in
2754 the Iberian Peninsula in 2023 (Copernicus Emergency Management Service, 2023b). The BA
2755 was 90% forest, mostly pine (*Pinus pinaster*, *Juntaex.es*, 2023). It was a wind-driven fire and
2756 the Canadian FWI indexes indicate all fine fuel was available to burn and extreme fire
2757 behaviour (FWI>50). The fire significantly impacted the nesting of protected bird species and
2758 rainfall shortly after the wildfire caused important runoff, erosion, and disruption of water
2759 supply to the local population (Armero, *Hoy*, 2023).

2760
2761 The two other significant fires in Spain happened in the Canary Islands, the Puntagorda
2762 wildfire (14th July 2023, La Palma, 32 km²) and the Arafo-Candelaria wildfire (15th August
2763 2023, Tenerife, 123 km²; Copernicus Emergency Management Service, 2023c). The latter
2764 spread for 9 days and 94% of its area was forest under conservation status, mostly of Canary
2765 pine (*Pinus canariensis*). The fire was exacerbated by local topography and mostly low to
2766 moderate severity (Educación Forestal, 2023b). Nonetheless, 26,000 people were evacuated,
2767 364 farms and 246 buildings (none residential) were affected, smoke impacts were substantial
2768 and damage was estimated at 80.4M€.

2769
2770 Like in Spain, winter shrubland burning was relevant in continental Portugal (~50 km² in
2771 February) but subsequent significant wildfire activity was restricted to two fires. The Sarzedas
2772 (66 km² ha) and Baiona (75 km²) wildfires started on the 5th of August under extreme fire
2773 weather (FWI>50), and burned mostly (~70%) forest, respectively of pine (*P. pinaster*) and
2774 eucalypt (*Eucalyptus globulus*) stands (Direção Nacional de Gestão do Programa de Fogos
2775 Rurais, 2023). Prevailing burn severity was moderate and damage to infrastructure and
2776 emergency restoration amounted to 6.4 M€ cost for the Sarzedas fire and a forest value loss
2777 of 1.4 M€ (Instituto da Conservação da Natureza e das Florestas, 2023). The major run of the
2778 Baiona wildfire was on August 7, corresponding to 73% of the total BA, when it threatened
2779 wildland-urban interfaces and damaged several buildings; one camping park and 20 small
2780 villages were evacuated (*Economia Online*, 2023). Moderate to high burn severity classes
2781 were dominant and costs were estimated at 2.7 M€ (tourism) and 7 M€ (houses), plus 1.4 M€
2782 in forest values loss and 2.9 M€ for emergency stabilisation (*Rádio e Televisão de Portugal*,
2783 2023; SIC Notícias, 2023). Finally, on 12th October, and under anomalously extreme fire
2784 weather for the time of the year, the Ponta do Pargo wildfire burned 48 km² in Madeira island,
2785 with an estimated agriculture-related cost of 3 M€ (*Rádio e Televisão de Portugal*, 2023).

2786
2787 The year was also mild in other European countries where burned areas can be extensive,
2788 namely Romania, Hungary, and Poland, which collectively summed only ~210 km² burned.
2789 EFFIS recorded 2461 km² burned in Ukraine, the largest fire attaining 42 km², but these figures
2790 are far from those registered in recent years. In northern Europe, a notable fire occurred in
2791 Scotland near Cannich, during the spring, the primary fire season in the humid Atlantic climate
2792 of the UK (Belcher et al., 2021). It started on 19th of April and burned ~33 km² of mainly
2793 moorland making it one of the largest fires in the UK in recent history (Sabljak, *The Herald*,
2794 2023; personal communication Niall MacLennan, Scottish Fire and Rescue Service).

2795
2796



2797
 2798 **Figure A1:** Summary of the 2023-2024 fire season in Greece. Time series of annual fire count,
 2799 BA, C emissions, PM2.5 emissions, 95th percentile fire size, fastest daily rate of growth, and
 2800 95th percentile fire daily rate of growth. Black dots show annual values prior to the latest fire
 2801 season, red dots the values during the latest fire season, and blue dashed lines the average
 2802 values across all fire seasons.
 2803

2804 7.4 North America

2805
 2806 Wildfire across North America in 2023-2024 was characterised by record fire activity across
 2807 Canada, lower than normal BA in the most flammable regions of the western US, near-average
 2808 fire activity across Mexico, and several extreme events that resulted in disastrous impacts to
 2809 human communities (Kolden et al., 2024). Over 150,000 km² burned in Canada in 2023
 2810 according to national statistics, over twice the previous record and over seven times the annual
 2811 average (Jain et al. 2024). The US burned 10,900 km² in 2023, well below the long-term
 2812 annual average (National Interagency Fire Center [NIFC], 2024a). Mexico has a relatively
 2813 short national wildfire recording system, but March 2023-February 2024 saw among the
 2814 highest area burned in the last decade (10,000 km²) and this has been associated with
 2815 ongoing drought conditions (Comisión Nacional Forestal, 2024).
 2816

2817 The fire season began earlier than normal in Canada, which regional experts have linked to
 2818 early snowmelt across much of the country and persistent drought conditions in the west.
 2819 Abnormally high temperatures and lack of rainfall also saw forested regions of eastern
 2820 Canada, including Quebec, transition rapidly to drought conditions at the end of May. British
 2821 Columbia province recorded its first wildfire evacuation in mid-April, and in late May, over
 2822 16,000 people were evacuated from Halifax, the capital city of Nova Scotia, which saw its
 2823 largest ever wildfire in 2023 (Jain et al. 2024). In June, two lightning outbreaks in Quebec
 2824 initiated several hundred new fires in what would eventually become a record BA year for the
 2825 province (4,300 km²) (Boulanger et al., 2024). While the majority of the Quebec fires were in

2826 remote regions, the smoke they generated was carried to several major cities in eastern North
2827 America, including New York which experienced its worst air quality in half a century as the
2828 observed daily mean PM_{2.5} concentration rose to 148.3 µg m⁻³, over 4 times the recommended
2829 daily limit (Wang et al. 2023). In total over 50 million people were exposed to high levels of
2830 PM_{2.5} for several days (Yu et al. 2024). This situation was further exacerbated in New York
2831 City by several wildfires in the nearby New Jersey pine barrens, a fire-prone dry pine forest
2832 that sees large fires during periods of drought.

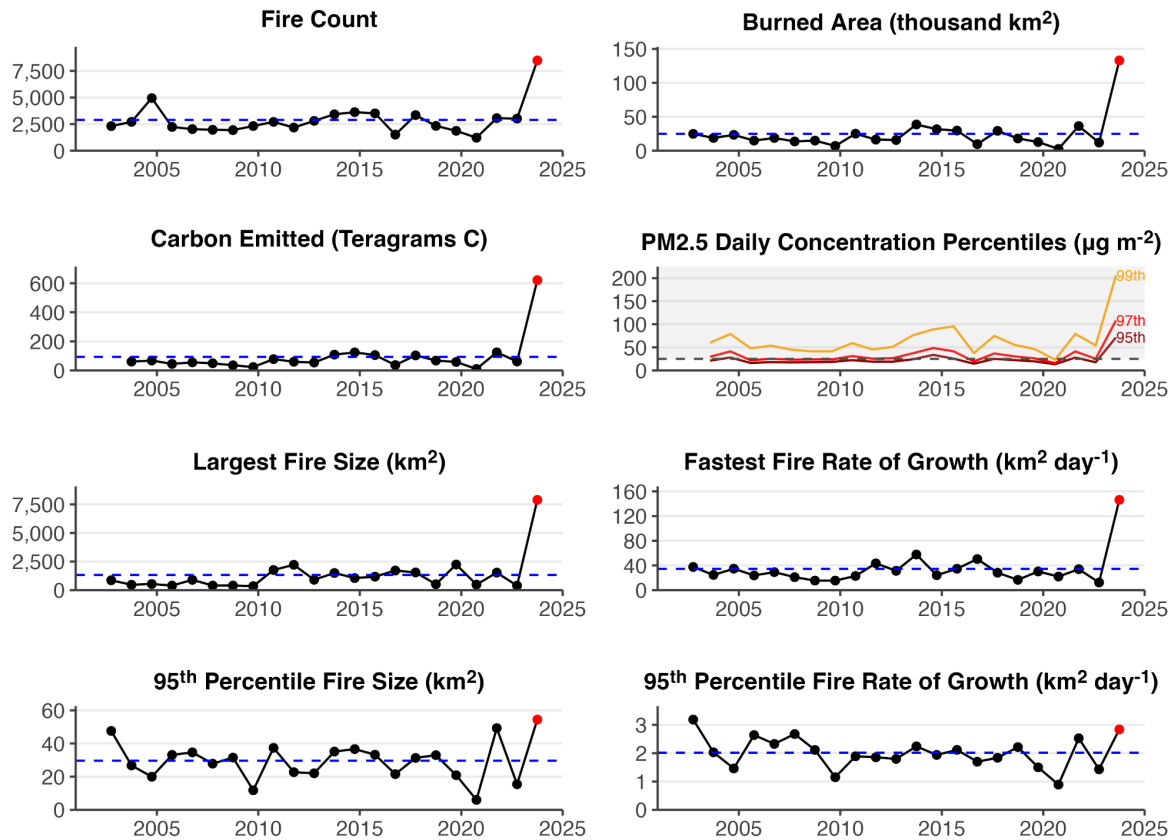
2833
2834 The year started with low fire activity across the USA. In the high plains of the central US, an
2835 outbreak of large wildfires occurred coincident with dry conditions and strong winds in March-
2836 April 2023. One wind-driven wildfire started by power lines in Oklahoma destroyed several
2837 dozen homes (Oklahoma Department of Emergency Management, 2023). Outside of the high
2838 plains, dry conditions also elevated fire activity across the Southern, Eastern, and New
2839 England regions of the USA. Mexico saw slightly above average fire activity during spring,
2840 which is the peak period of the fire season as debris burning and field clearing activities provide
2841 ignitions for predominantly shrubland and grassland wildfires.

2842
2843 As summer arrived in Canada, the western and boreal provinces and territories saw extreme
2844 and widespread fire activity, even as Quebec continued to burn. By the end of the year, record
2845 area had burned in British Columbia (2,300 km²), Alberta (2,700 km²), and the Northwest
2846 Territories (3,500 km²) accompanied evacuations of 232,000 people in numerous rural villages
2847 and large cities such as Yellowknife, NT and Kelowna/West Kelowna, BC, where a wildfire
2848 jumped the 2 km-wide Okanagan Lake (Jain et al., 2024; *CBC News*, 2023). The extreme
2849 behaviour of these fires not only shrouded large swaths of North America in smoke, but also
2850 generated an unprecedented 140 pyrocumulonimbus clouds (Jain et al. 2024). Eight
2851 firefighters were killed during summer 2023 in Canada (Jain et al. 2024), but miraculously no
2852 civilians died directly in the fires. Canada was at the highest National Preparedness Level 5
2853 for an unprecedented 120 continuous days starting on May 11, indicative of the significant
2854 resource sharing required by fire management; in all, over 5500 international personnel from
2855 12 countries and the EU were deployed to Canada during the 2023 fire season (Canadian
2856 Interagency Forest Fire Centre, 2023).

2857
2858 In the US, a relatively low activity fire season became deadly in August owing to unusual
2859 weather conditions facilitating extreme fire behaviour in multiple areas around the country. On
2860 8th August, a pressure gradient-induced katabatic wind event fanned several small wildfires in
2861 Hawaii, and 101 civilians died as the town of Lahaina on the island of Maui was consumed in
2862 the worst wildfire disaster in the US in a century (Pyne, 2017). Over 2,000 homes were
2863 destroyed and over 10,000 people were displaced as a result. Fires with extreme behaviour
2864 killed five additional civilians in the US states of Washington (2 fatalities), Louisiana (2
2865 fatalities) and California (C. Kolden, unpublished data). These extreme events stood in
2866 contrast to overall low fire activity and were notable for where they occurred. Washington does
2867 not typically see many extreme, wind-driven wildfires, and Louisiana is one of the wettest
2868 states in the US. By the end of August, the US BA was only 40 percent of normal and the
2869 lowest since at least 2000 (NIFC, 2023; National Oceanic and Atmospheric Administration
2870 [NOAA], 2023).

2871
2872 As North America transitioned to fall and then winter, Canada continued to burn nearly a month
2873 longer than normal, with the last large fires not controlled until late October. On 22nd
2874 September, remarkably late in the fire season, Canada saw its largest ever one-day total for
2875 BA at approximately 4,400 km² (Jain et al., 2024). While many of the Canadian fires were fully
2876 extinguished by winter, others simply smouldered into the deep peat layers, aided by a
2877 warmer-than-normal winter with a reduced snowpack. At the end of February 2024,
2878 spaceborne thermal sensors detected several dozen fires in northern Alberta and British

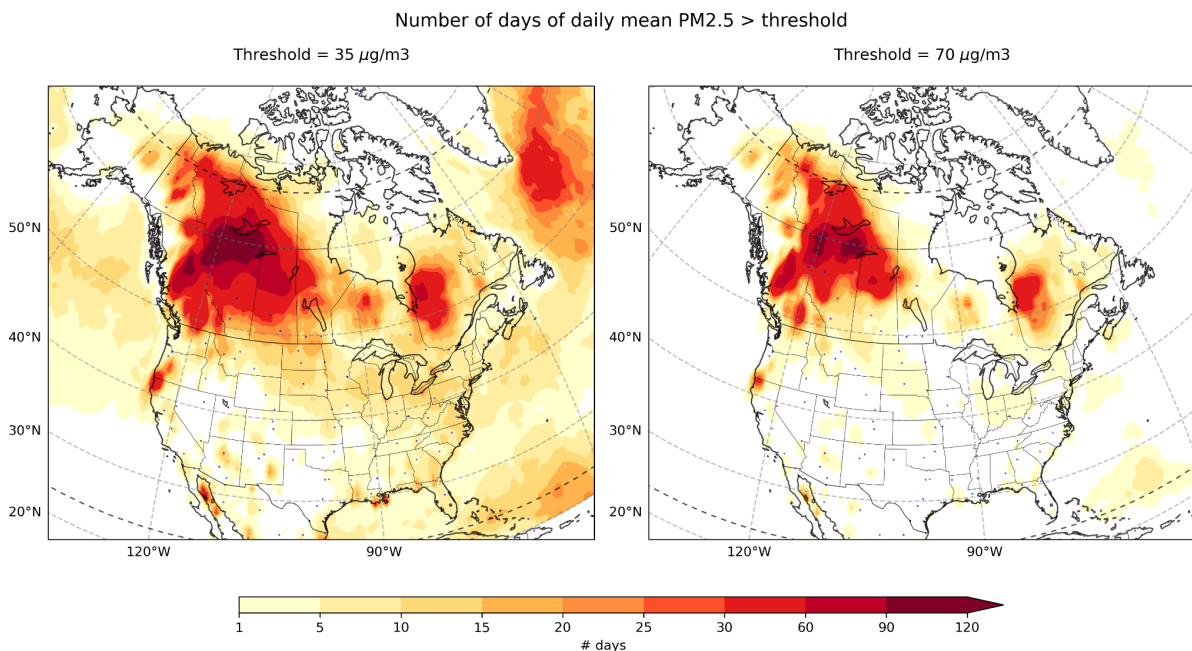
2879 Columbia that were overwintering fires, likely sustained by peat smouldering (Shingler, *CBC*
 2880 *News*, 2024; Scholten et al., 2021).
 2881



2882
 2883 **Figure A2:** Summary of the 2023-2024 fire season in Canada. Time series of annual fire
 2884 count, BA, C emissions, PM2.5 daily concentration percentiles (95th, 97th, and 99th), 95th
 2885 percentile fire size, fastest daily rate of growth, and 95th percentile fire daily rate of growth.
 2886 Black dots show annual values prior to the latest fire season, red dots the values during the
 2887 latest fire season, and blue dashed lines the average values across all fire seasons. The
 2888 PM2.5 daily concentration percentiles are based on area-weighted daily mean surface PM2.5
 2889 concentrations within each year across Canada from the CAMS atmospheric reanalysis
 2890 (Inness et al., 2019). The grey zone marks concentrations exceeding reference levels for 24-
 2891 hour mean PM2.5 concentration in Canada ($25 \mu\text{g m}^{-3}$) Canadian Environmental Protection
 2892 Act Federal-Provincial Working Group on Air Quality, 1998).
 2893
 2894
 2895

2896 US fire agencies recorded just over 10,900 km² burned in 2023, just over half of the 20-year
 2897 mean of 29,000 km² (NIFC, 2024a). Notably, over half of the BA was associated with higher
 2898 fire activity in the central plains grasslands and the southeastern US, while below normal fire
 2899 activity characterised California and the western US throughout 2023 as the region exited a
 2900 multi-year drought. However, the number of fires recorded was only slightly lower than
 2901 average. This quiet pattern broke in February 2024, however, when drought conditions from
 2902 the high plains region of the US down into north central Mexico coupled with strong winds to
 2903 produce massive, fast moving wildfires across multiple states on both sides of the US-Mexico
 2904 border. The US state of Texas recorded its largest ever single fire at over 4,000 km²
 2905 (Smokehouse Creek fire) in late February and early March that destroyed 130 homes across
 2906 the high plains region of the central US (NIFC, 2024b). Two civilians were killed by the flames

2907 in the relatively rural area dominated by ranching, over 10,000 head of cattle died, and
 2908 damages are estimated to be at least \$4.6 million (NOAA, 2024).
 2909
 2910



2911
 2912 **Figure A3:** Impact of Canadian wildfires visible in air quality metrics for North America. Panels
 2913 show the number of days in 2023 with mean PM2.5 concentration over a threshold of (**left**
 2914 **panel**) 35 $\mu\text{g m}^{-3}$ and (**right panel**) 70 $\mu\text{g m}^{-3}$. Both the National Ambient Air Quality Standards
 2915 (NAAQS) in the USA and the Canadian Ambient Air Quality Standards (CAAQS) have
 2916 exposure targets of 35 $\mu\text{g}/\text{m}^3$ on average within a single day.

2917

2918 The impact of North American fires on air quality is significant with half of the PM2.5 in America
 2919 suggested to originate from fires (O'Dell et al. 2019). An exceptional fire season, such as seen
 2920 in Canada in 2023, therefore poses an elevated level of health risk. Canada's wildfires
 2921 produced levels of PM2.5 across the country that were well in excess of the last 20 years
 2922 (**Figure A2**). Additionally, long-range transport of pollution from Canada affected the Pacific
 2923 Northwest, northern Midwest, and many eastern states (**Figure A3**). According to the National
 2924 Ambient Air Quality Standards (NAAQS) in the USA, the threshold for PM2.5 exposure is not
 2925 to exceed 35 $\mu\text{g}/\text{m}^3$ on average within a single day. CAMS analysis suggests that people living
 2926 within over half of US states experienced up to 2 weeks of exposure at or above this level. In
 2927 Canada, the safe limit for PM2.5 exposure, as defined by the Canadian Ambient Air Quality
 2928 Standards (CAAQS), is also 35 $\mu\text{g}/\text{m}^3$ over a 24-hour period. However, the situation was worse
 2929 in Canada due to closer proximity to fires, with many territories along the border experiencing
 2930 up to a month of degraded air quality that exceeded national recommended exposure limits,
 2931 British Columbia possibly facing twice the number of days at up to 2 months, and the Northern
 2932 Territories potentially with 3 to 4 months of exposure.

2933 The scale of the impact becomes particularly evident when comparing the number of days at
 2934 double the exceedance level (70 $\mu\text{g}/\text{m}^3$) due to short-range versus long-range transport of
 2935 pollutants. In Canada, where the pollution sources were more localised, the number of days
 2936 above this higher level remains substantial, ranging from a week along the border to months
 2937 still at the fire epicentres. In contrast, the USA, affected primarily by long-range transport from
 2938 Canada, experienced approximately half the number of such high-pollution days. It is also
 2939 important to consider the context of interannual variability in fire occurrence. Last year the

2940 USA experienced its lowest number of fires in 2 decades (**Figure 4**), so most of the pollution
2941 impacts came from Canada.

2942 7.5 Oceania

2943
2944 As is commonly the case, there was a marked latitudinal difference in wildfire patterns in
2945 Oceania in 2023-24. Fire activity was above average in the savannahs, grasslands and
2946 shrublands of tropical, subtropical and arid northern Australia. In contrast, fire activity in the
2947 southern states of Australia was generally below average, and well below the levels seen
2948 during the high impacting 'Black Summer' fires of 2019-20. In New Zealand and the Pacific
2949 Islands, fire activity was relatively low compared to the preceding two decades.

2950
2951 Given the vast scale of savannah fires, 2023-24 ranked among the top five years in BA for
2952 Australia as a whole since 2002 (Fisher, 2024; **Figure 2**). Fire in tropical and arid areas is
2953 tightly linked to rainfall in the preceding season (Alvarado et al., 2020). The above average
2954 fire seasons in the Northern Territory and northern Western Australia were driven to a large
2955 extent by elevated fuel growth associated with the La Nina conditions of the previous three
2956 years. These fires represented the vast majority of areas burned across the country in 2023-
2957 24 (**Figure S7**).

2958
2959 In the monsoonal north, savannah fires follow a strong seasonal pattern, with regular summer
2960 rain predictably followed by fire in the dry winter and spring months. In arid regions further
2961 south, fire remains tightly coupled to rain, but the seasonality is less pronounced. Anomalously
2962 large fires began as early as May and June in Western Australia and the Northern Territory
2963 respectively, continuing to as late as January.

2964
2965 The year was also marked by a series of early season, high impact fires in populated areas of
2966 southwestern Western Australia, southeast Queensland, NSW, Victoria and Tasmania. Hot,
2967 dry, windy conditions, and extended dry periods, are a major driver of forest, woodland and
2968 shrubland fires of the subtropical and temperate south of Australia (Collins et al., 2022). In
2969 addition to 2023 being the eighth-warmest year on record, the three months from August to
2970 October were the driest in over 100 years of records (Bureau of Meteorology, 2024).

2971
2972 From October to January a string of fire events led to loss of life, property and a range of other
2973 human and environmental impacts throughout the country's southeast and southwest. In some
2974 cases, significant fire activity was observed in areas impacted by the 2019-20 fire season.
2975 Despite these impacts, average rain in southern and eastern parts of Australia tempered fire
2976 activity for the austral summer. In Queensland and NSW, large fires in remote areas pushed
2977 the total BA in line with the long term mean, but this figure was well below average in Victoria,
2978 the Australian Capital Territory and Tasmania.

2979
2980 In the southwest of Australia, a volunteer firefighter was killed while responding to a fire near
2981 Esperance. The Kings Park fire in October occurred in a popular tourist area containing
2982 vulnerable flora and threatened Perth Children's Hospital. Perth was again affected by fires in
2983 December, with several injured and five homes lost. A further two homes were lost in the
2984 region in mid-January from fires that burned 60 km². A similar sized fire burned through rugged
2985 terrain in the Gammon Ranges 600 km north of Adelaide, threatening highly significant cultural
2986 and ecological values.

2987
2988 A large number of significant fires affected the eastern States of Australia in October. The
2989 Tara and Mount Isa fires in Queensland burned well over 400 km² combined, destroying 65
2990 homes and claiming the lives of two people. International and interstate support were deployed
2991 from New Zealand and Victoria to respond to the fires. Further south in New South Wales,
2992 significant fires included the Coolagolite Rd fire in Bega (over 70 km², 2 houses destroyed),
2993 the Willi Willi Rd fire in Kempsey (over 290 km², 8 houses destroyed, one person killed) and

2994 large fires around Tenterfield (approximately 300 km², four homes destroyed). In November
2995 the Hudson Fire burned 228 km², destroyed 4 properties and led to the death of a volunteer
2996 fire fighter, who was killed by a falling tree while fighting the fire. In December the Duck Creek
2997 Pilliga Forest fire burned 1,385 km² and initiated 3 documented fire-generated thunderstorms,
2998 with smoke impacts extending 500 km away and reaching Sydney.
2999

3000 In neighbouring Victoria the fire season was bookended by high impact events in October and
3001 then in February and March. Fires in Gippsland during October totalled 120 km² and exhibited
3002 some overnight fire runs that were regarded as unusual (Mills et al., 2022). An extended dry
3003 period saw fires impacting towns in central and western parts of the state in late February and
3004 early March. Over 40 homes were lost and five firefighters were injured fighting two fires that
3005 originated in the Grampians National Park and burned 60 km². Interactions with the
3006 atmosphere and topography were suggested to explain extreme behaviour that was reported.
3007 This fire was followed by another near Ballarat affecting grass, forest and a pine plantation.
3008 Despite several extreme fire weather days and evacuation advice, a significant suppression
3009 effort aided by interstate deployments minimised impacts. The fire burned over 200 km².
3010

3011 In the island state of Tasmania the fire season began early with the Coles Bay Bushfire burning
3012 27 km² of both private land and national park in September and then fires on Flinders Island
3013 in October. Other impactful fires that occurred during the season include the Dolphin Sands
3014 fire on the east coast of Tasmania that destroyed two homes and burned 2.5 km² and the
3015 multiple fires in the Brady lake area (Tasmania's central highlands) in February that destroyed
3016 two homes and burned up to 100 km² and a fire in the Waterhouse Conservation areas that
3017 required campers to evacuate.
3018

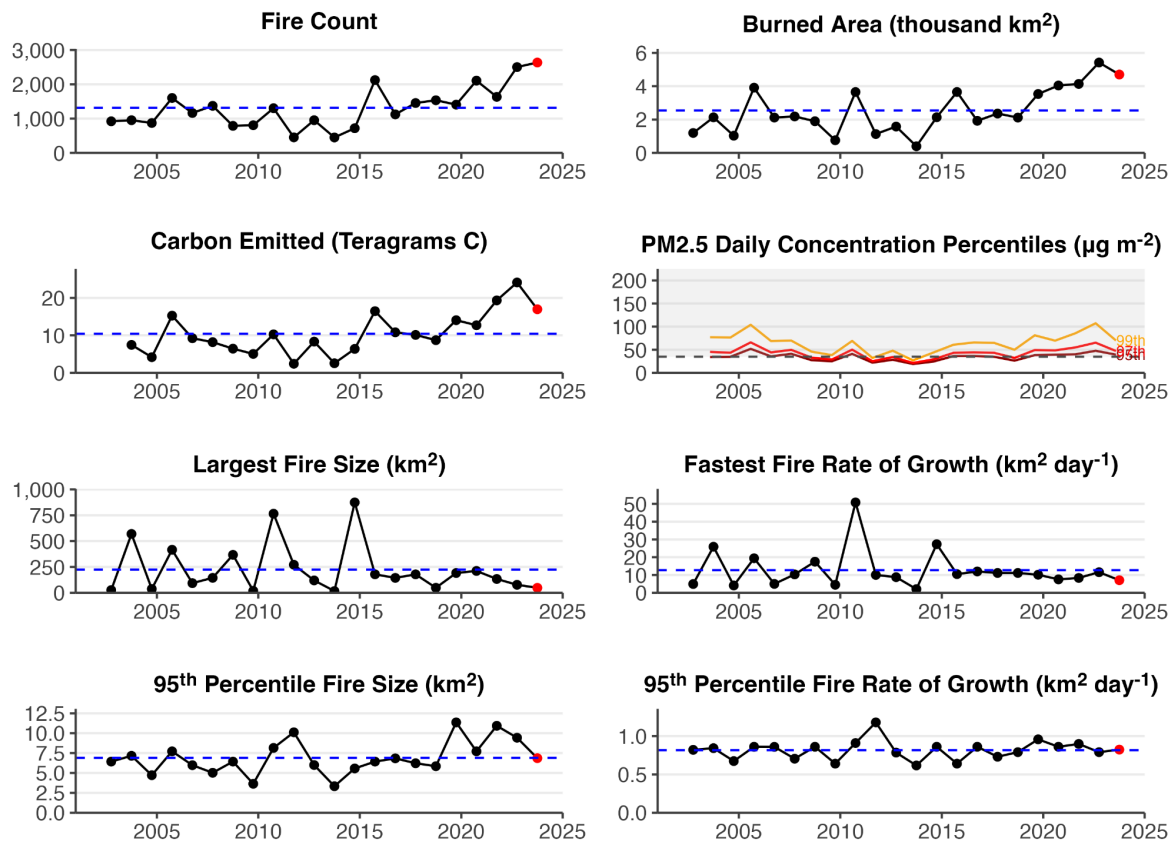
3019 New Zealand experienced a normal fire season after three well below average seasons prior
3020 under La Nina conditions. The fire season began early with a relatively large fire in September
3021 on the western side of Lake Pukaki in the central South Island in wilding pines. This fire totalled
3022 29 km² and this was the third major wildfire event in recent years in this area at an earlier than
3023 normal stage of the fire season, following the 2020 Pukaki (Aug.) and Ohau (Oct.) fires. New
3024 Zealand then experienced a spate of fires around Canterbury, on the South Island between
3025 late January and mid February 2024 with several houses burned and farmlands affected.
3026

3027 7.6 South America

3028

3029 The 2023-24 fire season in South America was characterised by a moderately below average
3030 fire activity but with positive wildfire anomalies in specific regions, which were reportedly
3031 exacerbated by extended periods of drought and heatwave across the continent (Clarke et al.,
3032 2024; **Figure 2, Figure 3, Figure 4**). In the Brazilian State of Amazonas, which features the
3033 largest extent of preserved old growth forests in Amazonia, June and October 2023 saw the
3034 highest fire counts since records began in 1998 (National Institute for Space Research, 2024;
3035 see also **Figure A4**). This continues a recent trend towards record-setting months for fire in
3036 Amazonas state, with new maxima being set in 7 months of the year since 2019 (National
3037 Institute for Space Research, 2024). Recent changes in deforestation and land use patterns
3038 are contributing to elevated fire ignitions in the state, reportedly compounded in 2023 by a
3039 historic drought and heatwave driven by El Niño (Espinoza et al., 2024, Clarke et al., 2024).
3040 Due to emissions of wildfire smoke, many areas of Amazonas experienced poor air quality
3041 from September to December 2023, including in the state capital, Manaus, where over 2
3042 million people were exposed to the second-worst air quality in the world in October (Ministério
3043 Público Federal, 2023). The event was so severe that, in November 2023, the Federal Public
3044 Ministry opened a Civil Action case against the State of Amazonas, demanding evidence that
3045 the State was investing in fire prevention and combat in line with the Plan for the Prevention
3046 and Control of Deforestation and Fires (Estado do Amazonas, 2020). This procedure
3047 evaluates whether Amazonas authorities are accountable for environmental damage causing

3048 severe air pollution, reflecting the Public Ministry's growing involvement at both federal and
 3049 state levels in monitoring environmental degradation and seeking to make authority figures
 3050 accountable (Ministério Público Federal, 2023).
 3051
 3052



3053
 3054 **Figure A4:** Summary of the 2023-2024 fire season in the Brazilian state of Amazonas. Time
 3055 series of annual fire count, BA, C emissions, PM2.5 emissions, 95th percentile fire size, fastest
 3056 daily rate of growth, and 95th percentile fire daily rate of growth. Black dots show annual
 3057 values prior to the latest fire season, red dots the values during the latest fire season, and blue
 3058 dashed lines the average values across all fire seasons. The PM2.5 daily concentration
 3059 percentiles are based on area-weighted daily mean surface PM2.5 concentrations within each
 3060 year across Amazonas from the CAMS atmospheric reanalysis (Inness et al., 2019). The grey
 3061 zone marks concentrations exceeding the reference level for 24-hour mean PM2.5
 3062 concentration in Brazil's National Air Quality Guidelines ($25 \mu\text{g}/\text{m}^3$; Siciliano et al., 2020).
 3063

3064
 3065 National fire monitoring systems in Brazil indicate that some areas of Amazonia experienced
 3066 anomalies in BA at the sub-state level. For example, BA in the municipality of Santarém in the
 3067 State of Pará rose from an average of 70 km^2 in 2019 to 2022 to over $1,000 \text{ km}^2$ in 2023, and
 3068 has already exceeded 250 km^2 in 2024 (MapBiomass Brasil, 2024). Similarly, in neighbouring
 3069 Belterra municipality, BA extent was more than 3 times greater during the year 2023 than in
 3070 2019-2022 (MapBiomass Brasil, 2024). In Floresta Nacional do Tapajós, one of the most
 3071 studied forest sites in the Amazon which spans Satarém and Belterra, forest fires accounted
 3072 for more than 60% of the burned areas (MapBiomass Brasil, 2024). 4 thousand people live in
 3073 24 communities of traditional and Indigenous populations in the region and depend on
 3074 protected forest resources for their cultural heritage, food security, economy and livelihood in
 3075 Floresta Nacional do Tapajós and Reserva Extrativista Tapajós-Arapiuns (Instituto Chico
 3076 Mendes de Conservação da Biodiversidade, 2019). Fires in 2023 compounded the challenges

3077 faced by these communities, who were already isolated by low river levels resulting from the
3078 drought that severely reduced their mobility and fishing, impacting on food security and
3079 enhancing socio-economic vulnerabilities.

3080
3081 In Chile, the 2023-24 fire season was marked by a significant escalation in both the number
3082 and size of wildfires, especially in the central and southern regions (**Figure 4**). Chile
3083 experienced its second-highest BA in the past 20 years (>4,000 km²; Jones et al., 2024). The
3084 peaks in BA at national scale were accompanied by peaks in the 95th percentile of fire size
3085 and daily rate of growth in highly-populated regions such as Valparaíso (**Figure 2, Figure 4**),
3086 indicating unusually large and fast-moving fires. These fires drew international attention due
3087 to their deadly impacts on society. In February 2024, severe wildfires struck the Valparaíso
3088 region in Chile, particularly affecting Viña del Mar and other surrounding areas (NASA Earth
3089 Observatory, 2024). These fires resulted in the deaths of at least 131 people and destroyed
3090 thousands of homes, leaving at least 1,600 people homeless (UN Resident Coordinator in
3091 Chile, 2024; *Al Jazeera*, 2024; *El Disconcierto*, 2024). The fires impacted the Lago Peñuelas
3092 National Reserve where more than 60 km² of forest were affected (Oberholtz, *Fox Weather*,
3093 2024). The National System for Disaster Prevention, Mitigation and Attention (SENAPRED)
3094 issued a red alert and ordered the evacuation of over 18 nearby towns (Oberholtz, *Fox*
3095 *Weather*, 2024). The February 2024 wildfires in Valparaíso followed other major disruptive
3096 wildfires in February 2023, which affected nearby regions of central Chile including Maule,
3097 Nuble, Bio bío, La Araucanía and Los Rios.

3098
3099 Several countries with territory in the west of Amazonia experienced anomalies in BA and fire
3100 behaviour during the 2023-24 fire season, which coincided with drought conditions. Peru's
3101 Loreto region, which neighbours the Brazilian state of Amazonas, faced its highest BA on
3102 record in the 2023-2024 fire season signalling the wider impacts of drought conditions in
3103 western Amazonia (**Figure 2**). The timing of peak anomalies in BA also coincided with those
3104 in Amazonas, around September-October 2023 (**Figure S2**). The northern Bolivian
3105 departments of La Paz and Beni experienced similarly-timed anomalies in BA. Remote parts
3106 of the Colombian Amazon also saw a significant uptick in BA since November 2023 which
3107 peaked in January 2024 (Mongabay, 2024). As a result of months of record-high temperatures
3108 and drought conditions since the beginning of El Niño, the region recorded higher C emissions,
3109 reflecting the severity of the burning at the end of the studied fire season. While the direct
3110 impacts on society throughout these regions was not as pointed as in the case of fires in Chile,
3111 the events are likely to have contributed to reductions in regional air quality and also impacted
3112 forest ecosystems and raised C emissions from fires in South America. At the end of January
3113 there was a wildfire in the mountains surrounding Bogota that affected the air quality that
3114 affected thousands of citizens of the capital of Colombia (France-Presse, *VOA News*, 2024).

3115
3116 In 2023-2024, Venezuela experienced its highest level of fire activity on record, particularly in
3117 January and February 2024 (ALER, 2024; *Tiempo*, 2024; **Figure 4, Figure S2, Figure S8**),
3118 notably affecting the states of Anzoátegui, Cojedes, Guárico, and Monagas, areas which
3119 dominant land cover primarily consists of savannas and extensive grasslands. This surge in
3120 fires during the dry season was intensified by unusually warm and dry conditions in the
3121 preceding months. These conditions, likely a result of global warming and changes in
3122 circulation and rainfall patterns associated with El Niño, making the landscapes more
3123 vulnerable to fires.

3124 **8 Appendix B: Frontiers in Observing and Modelling Extreme** 3125 **Fire Occurrence and Impact**

3126
3127 This appendix summarises current challenges in the observation and modelling of extreme
3128 fire events or seasons and identifies new technological and methodological advances that are
3129 raising opportunities to overcome such challenges. The section begins with a review of the

3130 study of extreme fire *occurrence* and then focuses on the study of extreme fire *impacts* across
3131 7 impact sectors. Advances in the study of both fire occurrence and fire impact are required
3132 because the impacts of extreme fires on society and the environment do not necessarily scale
3133 with observable fire properties (see **Section 6.2**).
3134

3135 **8.1 Occurrence**

3136

3137 **8.1.1 Definition of Extreme Fire Events**

3138

3139 Studying impactful fires or unusual fire seasons is crucial for understanding changes in
3140 exposure and vulnerability to fire. However, defining "extreme" events presents several
3141 challenges. These challenges can be categorised as follows.
3142

3143 Data-Oriented Challenges:

- 3144 ● **Lack of Consensus on Quantitative Criteria:** Variability in measurable criteria, such
3145 as fire size, across different studies hinders consistent classification. No statistical
3146 threshold is universally established to define outliers.
- 3147 ● **Geographic Variability:** Regional differences complicate universal definitions. Size
3148 thresholds vary widely, influenced by local fire regimes.
- 3149 ● **Evolving Definitions:** The term "extreme fire" has expanded over time, encompassing
3150 more fire types and behaviours. Climate change increases fire severity, suggesting
3151 definitions need flexibility.
- 3152 ● **Context Dependence:** Definitions vary with ecosystem types and fire histories,
3153 lacking standardisation on baselines such as fire return intervals or ecosystem
3154 damage.
3155

3156 Knowledge-Oriented Challenges:

- 3157 ● **Lack of Consensus on Qualitative Criteria:** Variability in criteria like fire behaviour
3158 and impacts reflects differing expert opinions. The subjective nature of significant
3159 impact complicates a clear definition.
- 3160 ● **Terminological Overlap and Redundancy:** Terms like "catastrophic fire" and "megafire"
3161 overlap, causing confusion due to unclear or interchangeable usage.
- 3162 ● **Influence of Language and Culture:** Interpretations differ across languages and
3163 cultures, affecting global reporting and definitions.
- 3164 ● **Societal Influence on Scientific Terminology:** Scientific terminology evolves with
3165 societal context. Language in scientific communication must remain adaptable and
3166 relevant to non-scientific audiences.
- 3167 ● **Scientific Rigour and Clarity:** Clear, consistent, and scientifically rigorous definitions
3168 are needed for standardised measurements. Existing definitions often fall short.
3169

3170 Defining 'extreme fire' requires a significant and inclusive effort across the fire science
3171 community. We avoided a strict definition here, instead adopting a broad and flexible definition
3172 as discussed in **Section 2.1**. To support a formal definition for future iterations of this report,
3173 standardised criteria and protocols should be developed through transdisciplinary approaches,
3174 such as through workshops involving input from scientists, fire practitioners, legislators, and
3175 communities (Chu et al., 2023; Linley et al., 2022; Shuman et al., 2022).
3176

3177 **8.1.2 Observation of Extreme Fire Events**

3178

3179 Global-scale data for characterising extreme fire events are primarily sourced from satellite
3180 observations, notably active fire detections, BA maps, and tracking of smoke plumes.
3181 However, to accurately define how extreme a fire event is, it is crucial to contextualise present-
3182 day observations within historical data. Unfortunately, the historical records of satellite-derived
3183 active fire and BA products are relatively short. The longest coherent observations on a global

3184 scale are derived from the MODIS instruments onboard the Aqua and Terra satellites,
3185 launched in 1999 and 2002 respectively. Various global BA products, such as MCD64 product
3186 family (Giglio et al., 2018) and FireCCI51 (Lizundia-Loiola et al., 2020), as well as active fire
3187 data like the MCD14 product family (Giglio et al., 2016), have been generated based on
3188 imagery acquired by MODIS. Although these time-series now span over two decades, they
3189 are still relatively short when compared to the decadal to centurial fire return intervals observed
3190 in many ecosystems.

3191
3192 Pre-MODIS satellite data, like that from the AVHRR program, exists and provides a continuous
3193 imagery archive from 1982 onwards. Although efforts are ongoing to generate a coherent BA
3194 product from AVHRR data (Otón et al., 2021), there are limits to the global applicability of
3195 these products. For example, unresolved challenges stemming from coherence issues
3196 between imagery from different AVHRR sensors result in artefacts and spurious trends in
3197 various regions worldwide (Giglio and Roy, 2022), although this has been debated by other
3198 authors (Pullabhotla et al., 2023).

3199
3200 Efforts are ongoing to extend the MODIS time-series by incorporating active fire data from
3201 ATSR and VIRS with BA data (e.g. Chen et al. 2023). However, due to the different
3202 characteristics of these data, creating a coherent, multi-satellite time-series of active fire data
3203 and/or BA is not straightforward. Concerns also arise with the impending decommissioning of
3204 MODIS-Terra, raising doubts about the continuity of existing long-term fire records. However,
3205 operational satellite sensors such as VIIRS onboard NOAA's series of satellites and SLSTR
3206 onboard the Sentinel-3 satellites offer promising capabilities for medium-resolution BA
3207 mapping (e.g. Román et al., 2024; Lizundia-Loiola et al., 2022). Urgent attention should be
3208 directed towards developing methodologies to integrate these new datasets into a coherent,
3209 long-term BA dataset. Furthermore, advancements in medium-resolution satellite data
3210 availability and revisit times, particularly from Landsat and Sentinel-2, now enable global BA
3211 mapping at spatial resolutions as fine as 20-30m (e.g., Roteta et al., 2021; Chuvieco et al.,
3212 2022), suggesting a potential future direction for coherent long-term global BA monitoring.

3213
3214 While BA is a key variable to characterise extreme fire occurrence, multiple other aspects of
3215 extreme fires can be characterised using satellite data. BA products can be used to cluster
3216 burned pixels in burned patches to obtain the number and size of individual fires (Archibald
3217 and Roy, 2009). Furthermore, the daily fire rate-of-spread, length of the active fire line and
3218 spread direction can be extracted from the daily fire expansion (Andela et al., 2019). These
3219 algorithms, such as the Global Fire Atlas used in this study, give global scale, coherent
3220 estimates of patterns and trends in fire number, fire size and rate of spread. However, these
3221 algorithms are sensitive to the temporal accuracy of the per-pixel burned date detection, the
3222 spatial resolution of the BA product, and any errors within each product. Recent advances
3223 focusing on clustering VIIRS thermal anomalies and extracting fire rate of spread, fire
3224 expansion and length of the active fire line show promising results (Andela et al., 2022,
3225 Hantson et al., 2022, Chen et al., 2022) but have so far not been developed globally. Future
3226 development towards a global product should allow for a more detailed characterization of fire
3227 characteristics in near real-time, well-suited for detection and quantification of fire extremes.

3228
3229 Active fire detections also record the amount of radiation emitted by the fire at the moment of
3230 satellite overpass (Fire Radiative Power, FRP), within the pixel detected by the satellite. While
3231 this information is related to the intensity of the fire, the usage of FRP has been difficult as a
3232 low-intensity fire burning a large extent of the pixel can have a higher FRP than a high-intensity
3233 fire burning a small fraction of a pixel. These complications have limited a more standardised
3234 and operational usage of FRP for quantifying fire extremes. Advances in active fire detection
3235 from higher resolution sensors may allow for a more comprehensive estimate of fire intensity
3236 when combined with FRP estimates from coarse resolution sensors (Schroeder et al., 2014,
3237 2016).

3238

3239 A natural starting point for this global assessment of the 2023-24 fire season was global data
3240 provided by the MODIS BA dataset, though we note that various national, state-level or
3241 regional systems exist and can add longer-term context to the extremity of fire seasons (e.g.
3242 Canadell et al., 2021; Short, 2014; Gincheva et al., 2024). Regional datasets generally depend
3243 on manual logging of fires via field approaches or desk-based identification with high-
3244 resolution imagery, or alternatively harness different blends of satellite observation with fire
3245 detection algorithms that can be regionally optimised. These approaches carry their own
3246 uncertainties and are limited by design to targeted regions, however their major advantage is
3247 multi-decadal coverage. Advances in compiling regional datasets into gridded records with
3248 global coverage are bringing these advantages to formats compatible with global scale
3249 analysis (Gincheva et al. 2024) and will thus be explored in future efforts to characterise
3250 regional extremes at scale.

3251

3252 **8.1.3 Prediction of Extreme Fire Events**

3253

3254 Since the 1970s, fire predictions have relied on empirical fire behaviour models tailored to
3255 specific ecosystems (Bradshaw et al., 1983; Noble et al., 1980; Stocks et al., 1980; van
3256 Wagner, 1987), becoming pivotal tools for fire management agencies (San-Miguel-Ayanz et
3257 al., 2013). The ease of implementation and the availability of weather data have contributed
3258 significantly to their widespread adoption. However, despite their utility, several studies have
3259 highlighted the limited effectiveness of the FWI and similar metrics in fuel-limited ecosystems,
3260 where fires are driven by the short-term superficial drying of intermittently available biomass
3261 (Yebra et al., 2013). The absence of consideration for actual fuel availability presents a
3262 constraint to the meaningful application of the FWI in savanna-type ecosystems. Likewise, the
3263 response of fuel moisture to meteorological factors can be influenced by external factors that
3264 are challenging to observe and model at scale, such as mortality triggered by insect infestation
3265 or disease (Canelles et al., 2021). Beyond weather conditions, the remaining prerequisites for
3266 fire activity—namely, fuel and ignitions—are intricately linked to vegetation state, lightning
3267 activity, and human behaviours. Improving fire forecasts beyond solely considering fire
3268 weather could be achievable by accurately describing these components. This has been
3269 widely recognised in the global vegetation-fire community for several decades (Hantson et al.,
3270 2016), and consequently great advances have been made to address this through the
3271 development of fire-enabled DGVMs, as used in this report. However, explicit representation
3272 of these processes introduces biases and instabilities that, when used in isolation, limits their
3273 utilage for assessing climate and human drivers of BA extremes (Hantson et al., 2020; Burton,
3274 Lampe et al., 2023).

3275

3276 The availability of remote observations for fuel, either independently (Yebra et al., 2018) or
3277 supported by modelling frameworks (McNorton and Di Giuseppe, 2024), has demonstrated
3278 potential in aiding the development of new fire models and indices that partially incorporate
3279 fuel considerations into their formulation (Di Giuseppe, 2023, Hantson et al., 2016). However,
3280 it is the emergence of the data-driven revolution that holds the promise of significantly
3281 enhancing our predictive capabilities for extreme fires (McNorton et al., 2024). This has driven
3282 the development of semi-empirical tools at regional and global scales that could improve fire
3283 predictions, and their effectiveness will be assessed in the next edition of the report. Not only
3284 can these tools enhance predictive capability for extreme fires, they also present an
3285 opportunity to disentangle the drivers of the prediction, giving us the capability to address or
3286 at least understand the causes of the event, as demonstrated by the POF and ConFire
3287 frameworks used here. The coupling of FireMIP models with observational data (Burton,
3288 Lampe et al., 2023 and used here), also showcases the potential to bridge the advanced
3289 modelling capabilities of FireMIP with application-specific approaches such as ConFire and
3290 POF.

3291

3292 Despite these technological advancements, widespread adoption is unlikely to occur
3293 suddenly, as there typically exists a delay between the creation of new indices, their

3294 operationalization, and operational implementation by those responsible for fire prevention
3295 and control. There are also likely to be some stubborn issues with the detail provided large-
3296 scale observational data available to predictive systems, particularly in the case of fuel loads
3297 and fuel state (e.g. living versus dead). New global biomass observations, such as those from
3298 airborne and spaceborne Light Detection and Ranging (Lidar) and Synthetic Aperture Radar
3299 (SAR), provide insights into fuel loading but they are not currently providing information
3300 regarding fuel state that would be useful for modelling fuel moisture response to
3301 meteorological conditions (Santoro et al., 2022; Hunke et al., 2023).

3302
3303 Another emerging element is the recent availability of fire danger predictive systems at the
3304 seasonal and subseasonal timescales (Di Giuseppe et al., 2024). Currently, there is limited
3305 evidence on how these longer-range tools could contribute to prevention planning and
3306 adaptation strategies. While they exhibit minimal skill beyond two months, they may offer
3307 valuable pre-seasonal warnings under specific conditions established during important
3308 atmospheric modes of variability.

3309
3310 The current report does not include hybrid models for seasonal fire risk, fire propagation
3311 models, or fire susceptibility/risk mapping tools. Incorporating these approaches could offer
3312 valuable insights and will be considered in future reports. These advanced models and tools,
3313 which account for both past and present weather conditions as well as other critical factors
3314 such as soil moisture and vegetation dryness, can enhance our understanding of fire dynamics
3315 and improve predictive capabilities. By exploring these methods, future editions of the report
3316 could provide a more comprehensive overview of fire risk assessment and management
3317 strategies.

3318

3319 **8.1.4 Attribution Extreme Fires to Global Change**

3320

3321 The prediction and management of extreme fire events have become increasingly complex
3322 due to the multifaceted impacts of global change. Climate change exacerbates fire risks
3323 through rising temperatures, altered precipitation patterns, and more frequent and severe
3324 droughts, as shown in Canada and western Amazonia in this report. These climatic shifts
3325 affect vegetation productivity, with elevated CO₂ levels potentially increasing biomass and
3326 thereby providing more fuel for fires. Nutrient deposition and other environmental changes
3327 influence ecosystem responses, further altering fire potential. Land use changes and
3328 management practices also significantly influence fire dynamics. For example, human
3329 activities such as deforestation, urban expansion, and agricultural practices can both mitigate
3330 and exacerbate fire risks, with socioeconomic factors shown to have a strong influence on
3331 overall extreme fire likelihood in western Amazonia, and potentially contributing to increases
3332 in BA in 2023 in some areas of Greece. Effective land management strategies, including
3333 prescribed burns and forest thinning, are crucial for reducing fuel loads and minimising fire
3334 impacts. While climate driven estimates of extreme behaviour are plentiful, few modelling
3335 frameworks take into account most of these dynamic factors and their interactions (Rabin et
3336 al., 2017).

3337

3338 We have used model-data fusion techniques that account for these factors in this report, and
3339 have been able to attribute some of their influences in certain places. This report utilises semi-
3340 empirical models that blend empirical data with process-based understanding to better predict
3341 fire behaviour. Quantifying uncertainty in these models is essential, especially when dealing
3342 with extremes. By generating probability distributions, researchers can better understand the
3343 likelihood of various fire scenarios, informing more effective management and policy
3344 decisions.

3345

3346 Through uncertainty quantification techniques, we have been able to ascertain where we are
3347 confident in our attributions. However, uncertainties still remain, many from not considering
3348 the complex interactions and feedback onto fire, some from fire itself as it consumes fuel, and

3349 affects from weather. Coupled vegetation-fire models explicitly represent many of these
3350 feedbacks. However, current FireMIP models struggle to accurately simulate extreme fire
3351 events (Hantson et al., 2020). One key factor hampering improvements in model development
3352 is our limited understanding of factors driving fire extinguishers in a natural setting. While much
3353 process based knowledge exists on the factors influencing fire start and fire spread, only
3354 limited knowledge exists on the myriad of factors that can stop a fire, from changes in fuel
3355 moisture, structure and heterogeneity to landscape fragmentation and fire fighting, and how
3356 these interact (e.g. Finney et al., 2012). Without a strong theoretical understanding of these
3357 factors, process based modelling of extremes at a global scale might be limited in the near
3358 future. For the 2023 focal events, we have shown that low fuel loads and variations in human
3359 modification of the landscape can limit fuel spread (**Figure 11, Figure 12, Figure S13, Figure**
3360 **S14**). However, we only look at a handful of events and further examples are required at larger
3361 scales to inform improvement of process-based rates of spread in fire models.

3362
3363 To move forward, we need to combine these concepts in attribution techniques and quantifying
3364 uncertainty with coupled vegetation fire models, such as in FireMIP. Early attempts of this are
3365 promising – ConFire (used in **Section 3, 4 and 5**) borrows many of the modelling concepts
3366 from FireMIP, though it still lacks many feedbacks from fire itself. We have also used the latest
3367 FireMIP models coupled to an uncertainty framework for broad-scale, uncertainty-based
3368 attribution (obtained from Burton, Lampe et al., 2023). But they struggle at reproducing the
3369 tails of distributions where extreme events are found. Another way to develop these
3370 techniques is to move towards an integrated system that would inform both attribution and
3371 future projections in a seamless way. We make some progress in this direction here using
3372 tools such as ConFire, by using the information gained from fire drivers to build future
3373 projections, however there is more work to do to link statistical approaches for today’s fires to
3374 future projections.

3375
3376 The human role in driving fire and extremes is hard to represent. Despite the often reported
3377 influences people have on both increasing extreme burning or causing the observed decline
3378 in global BA, the role of humans in the landscape remains hard to capture and on the whole,
3379 remains one of the most uncertain aspects of this report. Agent-based modelling (ABM) is
3380 trying to address this by simulating the behaviours and interactions of individual human entities
3381 (e.g., deforestation, crop residue burning, suppression, etc.) within a given environment (Ford
3382 et al., 2021). This approach provides a dynamic representation of how different factors
3383 contribute to fire risk and links well with subsequent sections of the report. These approaches
3384 could be a major contributor to subsequent issues of this report. However, the integration of
3385 these advanced modelling techniques into operational use faces challenges, as there is often
3386 a delay between the development of new approaches and their widespread adoption by fire
3387 management agencies. This underscores the need for continuous improvement and adoption
3388 of innovative modelling approaches to address the growing threat of extreme fire events
3389 effectively.

3390
3391 In addition to fire weather index (FWI) and BA projections, it is crucial to go beyond these
3392 metrics to consider wider impacts such as intensity, and emissions. Understanding the
3393 intensity of fires helps in assessing their destructive potential and the severity of their
3394 ecological and societal impacts. Emissions from fires, including C dioxide and other
3395 greenhouse gases, contribute to climate change and air quality issues. Finally, evaluating the
3396 broader impacts of fires, such as on biodiversity, human health, and economic stability, is
3397 essential for developing comprehensive adaptation and mitigation strategies. Quantifying and
3398 understanding the uncertainty in these projections is crucial for developing adaptive strategies
3399 that can effectively respond to the evolving fire risks posed by global change.

3400 3401 **8.1.5 Projection of Fire Extremes** 3402

3403 Projections of extreme fire events under future climate scenarios indicate a significant increase
3404 in their frequency and intensity. Semi-empirical models used in this report project that extreme
3405 BA events, currently rare, are likely to become more common by the end of the century. These
3406 projections highlight the urgent need for robust fire management strategies and policies to
3407 mitigate the impacts of these increasingly severe fire events on ecosystems, communities,
3408 and global C dynamics. Quantifying and understanding the uncertainty in these projections is
3409 crucial for developing adaptive strategies that can effectively respond to the evolving fire risks
3410 posed by global change. In our ConFire uncertainty quantification framework, we have been
3411 able to make some confident inferences about the potential state of wildfires in the coming
3412 decades. However, we have also identified that there is a significant amount of crucial
3413 information that is currently beyond our reach due to the uncertainties involved. Our ability to
3414 forecast for the upcoming season, as well as for the next 2-3 decades, requires further
3415 refinement as we are observing mixed and uncertain responses. Beyond that, we still show
3416 similar uncertainties in responses of Canada and Western Amazonia under different scenarios
3417 as highlighted by UNEP (2022a), which uses the previous generation of climate models from
3418 CMIP5 – shows a large overlap in the potential range of changes in the occurrence of fire
3419 extremes between SSP370 and SSP585. This does not imply that mitigation efforts for one
3420 scenario will be ineffective compared to another, but rather indicates a lack of understanding
3421 regarding the response of extremes to these scenarios. By narrowing down the uncertainty
3422 ranges, we can better target adaptation efforts and evaluate the effectiveness of mitigation
3423 strategies. The reduced likelihood of extreme event recurrence in our high mitigation, however,
3424 does show that we can start separating out how mitigation efforts might affect fire extremism,
3425 though not in the level of detail needed for policy.

3426
3427 There are three main ways we may be able to constrain uncertainties in the coming reports.
3428 The first is development of the underlying GCMs that project future change in the drivers of
3429 BA. For individual models, this is a slow process and, beyond informing CMIP model
3430 development, is outside what the State of Wildfires can contribute to. However, bringing in
3431 more models, including having another model to incorporate any remaining biases in
3432 simulated fire from the correct models into our uncertainty projection, will help us constrain
3433 uncertainties more (Kelley et al., 2023). The second is obtaining more information and
3434 understanding of how fire drivers relate to fire extremes as outlined in the previous section.

3435
3436 Better ways of describing the statistical relationship between observed and modelled climate,
3437 land surface and fire today is a third approach. Investigating the dynamical climatic drivers of
3438 extreme fire conditions in different regions can help to physically disentangle and potentially
3439 constrain sources of uncertainty in future climate projections, for example by constructing
3440 physical storylines (Shepherd et al., 2018; Mindlin et al., 2023). These storylines of plausible
3441 future change, or other similar approaches to quantify and explain uncertainty in projections,
3442 provide critical information for communities to develop robust adaptation strategies (Lemos et
3443 al., 2012) and prepare for future losses and damage caused by evolving fire risks posed by
3444 global change. Next to understanding future uncertainty, further insights into these dynamical
3445 drivers can support the development of improved physics-informed bias adjustment of climate
3446 models (Maraun et al. 2017). Currently available methods to bias adjust climate models for
3447 their use in fire models, such as the ISIMIP3BASD method, have been shown to modify the
3448 climate change trend, particularly in extreme threshold indices (Casanueva et al 2020, Spuler
3449 et al 2024) or increase spread in climate model projections (Lafferty and Sriver, 2023). Bias
3450 adjustment methods should therefore be evaluated carefully and leave scope for future
3451 method development that physically links present-day biases to future uncertainty.

3452
3453 **8.2 Impact**
3454

3455 **8.2.1 Direct Exposure of People and the Built Environment**

3456 The direct exposure of people and the built environment could be studied at scale, however
3457 these analyses have not been performed routinely and have thus far tended to focus on
3458 specific regions. The wildland urban interface (WUI) has been the focus of direct fire exposure
3459 to populations, particularly in urban conflagrations like the Lahaina fire in August 2023. Efforts
3460 to reduce fire losses in the WUI through prevention, fuel reduction, and mitigation (Calkin et
3461 al., 2023) are crucial due to increased populations in these areas (Radeloff et al., 2018) and
3462 rising fire potential, accelerating community impacts (Higuera et al., 2023). US studies showed
3463 a doubling of direct population exposure to large fires during 2000-2019, mainly due to fires
3464 encroaching on the WUI (Modaresi Rad et al., 2023). Similar increases were noted in wildfires
3465 affecting roads and energy infrastructure, complicating transportation and energy reliability.
3466 Most structure losses in the US occurred in grasslands and shrublands, not forests,
3467 highlighting the need to look beyond traditional forest-centric assessments. Global WUI
3468 mapping efforts (Schug et al., 2023) offer opportunities to identify vulnerable areas and
3469 characterise fire exposure trends (Chen et al., 2024; Tang et al., 2024).

3470 Understanding fire characteristics that result in direct exposure, structure loss, and fatalities is
3471 essential. For instance, Abatzoglou et al. (2023) found that fires driven by strong downslope
3472 winds in the western US caused most structure losses and fatalities from 1999-2020, despite
3473 being only 12% of all fires. These winds push fires downhill into WUI communities,
3474 overwhelming suppression efforts and fuel treatments. The 2023 fires in Hawaii, Greece, and
3475 Chile were driven by such conditions (Synolakis and Karagiannis, 2024). Diagnosing
3476 characteristics of extreme fires, including meteorological conditions (Van Wagtendonk, 2006;
3477 Lareau et al., 2018) and pre-existing vegetation conditions (Stephens et al., 2022), provides
3478 insights into fires likely to cause significant impacts, helping prioritise mitigation efforts.

3479 A significant gap in characterising extreme fires and their human impacts (fatalities,
3480 evacuations, structure loss, secondary morbidity, economic losses, and impacts on food,
3481 water, energy, and transportation) is the lack of comprehensive national-to-global data.
3482 Wildfire morbidity data are collected in few countries due to the infrequency of fatal wildfires
3483 and unclear cataloguing responsibilities (Haynes et al., 2019). Smoke-induced morbidity
3484 estimates rely on spikes in hospital visits linked to specific events (Johnston et al., 2021).
3485 California pioneered systematic cataloguing of structure losses in 2013, yielding significant
3486 insights (Kolden and Henson, 2019; Syphard and Keeley, 2019), but this model has not been
3487 widely adopted. Canada now catalogues wildfire evacuations, but complexities remain in
3488 characterising these events (Beverly and Bothell, 2011). Global insurance records document
3489 insured losses but do not represent broader losses due to high rates of uninsured property
3490 (Hazra and Gallagher, 2022).

3491 Databases like EM-DAT have attempted to fill this gap but often overgeneralize wildfire
3492 impacts, relying on variable accuracy news reporting. Expanding and improving quantification
3493 of wildfire impacts on humans is critical to overcoming the "burned area fallacy" and
3494 developing effective mitigation models (Kolden, 2020). Remote sensing for documenting
3495 structure loss and fire incursion into the WUI, combined with high-resolution sensors and air
3496 quality monitors, can facilitate interdisciplinary research on wildfire smoke and medical
3497 morbidity (Liang et al., 2021).

3498 **8.2.2 Air Quality and Health Impacts**

3501 The impacts of fire on air quality and health could be studied routinely at scale using
3502 atmospheric models, however these analyses face several challenges. Exposure to outdoor
3503 pollution is a major global health risk (Murray et al., 2020). Fine particles with a diameter less
3504 than 2.5 µm (PM2.5) are particularly concerning due to their link to cardiovascular diseases.

3505 Fire smoke is increasingly impacting air quality and is considered more toxic per unit of PM
3506 exposure than other pollution sources (Aguilera et al., 2021). The World Health Organization
3507 (WHO) has reduced the annual mean exposure limit for PM_{2.5} from 10 µg m⁻³ to 5 µg m⁻³.
3508 With 95% of the world's population exposed to PM_{2.5} concentrations of at least 10 µg m⁻³
3509 (Shaddick et al., 2018), the rise in severe wildfire pollution poses an elevated health risk.
3510

3511 Issues contributing to the challenge of quantifying the impact of fire pollution on human health
3512 are the same as those for other pollution sources, including, but not limited to, a lack of ground-
3513 based measurements in many regions of the world, a need for more pollution dispersion and
3514 transport studies, a deeper understanding of plume dynamics and chemistry, and a partial
3515 reliance on animal-based human exposure/response models (e.g., Fiore et al. 2012, Fuzzi et
3516 al. 2015). These issues contribute to three of the major challenges in quantifying the impact
3517 of extreme fire pollution on human health. The first is accurately measuring the amount of
3518 pollution that a wide variety of communities are exposed to and then attributing the contribution
3519 of a wildfire event, that could be 100s or 1000s of km away, to the measured concentration.
3520 The second is that PM_{2.5} is not the only pollutant of concern (the EPA regulates six pollutants
3521 of concern for American citizens and a wildfire produces them all). The third is accurately
3522 linking exposure to a wide range of pollutants to their associated short-term and long-term
3523 health impacts.
3524

3525 Tools to assess air quality primarily consist of ground-based measurements and modelling.
3526 Ground based observations provide an accurate measurement of pollution at their location.
3527 However, measurement locations are spatially sparse. Ongoing efforts to increase spatial
3528 coverage include deployment of small relatively affordable particulate matter sensors, such as
3529 the PurpleAir network, by a wide range of communities (not just scientists), and efforts to relate
3530 surface PM concentrations to measured aerosol optical depth from satellites (Li et al., 2021).
3531 One additional constraint of observations is that they cannot differentiate pollution sources,
3532 but modelling can.
3533

3534 Dispersion modelling uses emission estimates, reanalysis meteorology, and topography to
3535 provide estimates of ambient pollutant concentrations at varying spatiotemporal scales. A
3536 challenge for models currently lies in the large uncertainty in fire emissions (Reddington et al.,
3537 2016; Carter et al., 2020; Pan et al., 2020). Emission data will therefore require calibration
3538 against observations and adjusting before the contribution of fire to pollution can be quantified.
3539 Improved emission datasets will increase confidence in these assessments.
3540

3541 Environmental and personal factors both influence cardiovascular health, making it
3542 challenging to isolate the effects of fire smoke. Impacts may also not be immediate; some
3543 effects can be acute, such as exacerbation of asthma, while others emerge over a longer
3544 period, like the development of cardiovascular disease, and are thus much more difficult to
3545 directly connect to a specific extreme fire event. Conducting epidemiological studies that link
3546 fire smoke exposure to specific health outcomes requires comprehensive data collection and
3547 follow-up. These studies are resource-intensive, time consuming, and subject to potential
3548 limitations in data.
3549

3550 **8.2.3 Impacts on Indigenous and Traditional Communities**

3551 The impacts of fire on Indigenous and Traditional Communities are not studied routinely and
3552 at scale, typically focussing on isolated regions and specific communities. Indigenous peoples
3553 and local communities (IP&LC) are disproportionately exposed to extreme fire impacts
3554 because of their proximity to the land and resources from which their cultures, livelihoods and
3555 often food and medicines derive. Once landscapes are degraded through fire the access and
3556 abundance of various resources can be shifted. At the same time these communities are often
3557 less supported by the state due to access and their political and economic marginalisation
3558 linked to systemic socioeconomic disadvantages. These communities suffer not only post-fire
3559

3560 impacts, but can be disincentivized from particular land and resource use activities because
3561 of the increasing threat of forthcoming wildfires. Whilst the multiple important values (e.g.
3562 instrumental, intrinsic and relational) associated with landscapes by IP&LCs are threatened
3563 by fire, there is a lack of systematic pre- and post-fire assessment of these impacts (van
3564 Leeuwen and Miller-Sabbioni, 2023).

3565
3566 Historically fire governance has added an additional burden to IP&LC, often prohibiting cultural
3567 fire use and management to the detriment of local knowledge and values, and in some cases
3568 also increasing the propensity of wildfire in tropical systems as well as savannahs (Carmenta
3569 et al., 2019; Daeli et al., 2021; Croker et al., 2023). In some contexts there is a shift towards
3570 correcting these issues and renewed interest in and support for cultural burning and
3571 Indigenous approaches to land management. For example, integrated fire management is
3572 gaining traction globally and sits at the heart of a number of interventions and international
3573 policy efforts (e.g. Fire Hub, 2023). The premise of Integrated Fire Management (IFM) is to
3574 maximise the 'good' fire and minimise the occurrence of wildfire often through an approach of
3575 connecting knowledges (i.e. expert and place-based or Indigenous). Whilst nations sit at
3576 different stages of development in respect to IFM, and many IP&LC feel there is a long way to
3577 go, the growing interest is promising (Bilbao et al., 2019; Luque et al., 2020; Rodríguez-Trejo
3578 et al., 2022). Research is needed to better understand the effectiveness of IFM and what
3579 mechanisms and processes work best for rebalancing the influence of various forms of
3580 knowledge on fire management. For instance in North America where fire use is considered a
3581 form of medicine to the land, and anointed to particular patches of the landscape for care
3582 (Palmer, 2021). These approaches are perceived as potentially more just as they allow for the
3583 many meanings and uses of fire to exist and persist. For example, In Australia, the term
3584 'Country' is used to convey the cultural and spiritual connection of Aboriginal peoples to the
3585 land and water in specific regions. This link profoundly colours Indigenous peoples' experience
3586 of extreme fire events such as the Black Summer fires of 2019-20 (Nolan et al., 2021b).

3587 3588 **8.2.4 Economic Impacts**

3589
3590 The economic impacts of fire are generally not studied routinely and at scale, but rather focus
3591 on individual regions and fire seasons and arrive some time after the event occurs. Extreme
3592 wildfires cause economic disturbances worldwide, with impacts varying across regions due to
3593 different economic structures, environmental conditions, and response strategies. These
3594 impacts include property and infrastructure loss, business downtimes, supply chain
3595 disruptions, decreased tourism, health costs, reduced productivity, and damages to
3596 ecosystem services. While tangible costs (e.g., insured property losses) are easier to
3597 measure, intangible costs (e.g., lives lost, health impacts from smoke exposure, damage to
3598 species and habitats) are harder to quantify due to data availability and varying temporal and
3599 spatial scales. Consequently, assessing the true economic costs of extreme wildfires is
3600 challenging. Additionally, some sectors may benefit economically from post-fire reconstruction
3601 and suppression efforts (Nielsen-Pincus et al., 2014; Meier et al., 2023a).

3602
3603 Research on the economic impacts of wildfires has mainly focused on developed economies
3604 in Europe, the US, and Australia. For instance, Meier et al. (2023b) estimated economic losses
3605 from a 1-in-10 year extreme wildfire in Mediterranean Europe: €162–439 million in Portugal,
3606 €81–219 million in Spain, €41–290 million in Greece, and €18–78 million in Italy. California's
3607 2018 and 2020 wildfire seasons resulted in approximately \$150 billion and \$19 billion in
3608 economic damages, respectively (Wang et al., 2021; Safford et al., 2022). Australia's 2019-
3609 2020 fire season was the costliest natural disaster in the country's history, with an estimated
3610 GDP decrease of \$10 billion (Wittwer and Waschik, 2021).

3611
3612 While satellite data can provide proxies for economic impact shortly after events, traditional
3613 economic indicators such as sectoral GDP, employment, hospitalizations, suppression
3614 spending, house prices, and tourism revenues are often not publicly available, not

3615 harmonised, or unable to capture long-term effects. Thus, the full economic costs may only
3616 become apparent years later.

3617
3618 Econometric analysis of wildfires is more challenging than for other natural hazards because
3619 wildfire occurrence is largely influenced by human activity, land-use choices, and
3620 socioeconomic factors. While earthquakes are non-human-induced, making causal analysis
3621 easier, wildfires correlate with many economic outcome variables. Despite these challenges,
3622 counterfactual analyses or econometric approaches like instrumental variables can provide
3623 reliable causal estimates of wildfire impacts. These methods promise to enhance
3624 understanding and mitigation of wildfire economic consequences through more accurate
3625 analyses, helping target suppression policies and allocate resources effectively.

3626
3627

3628 **8.2.5 Loss of Biodiversity, Ecosystem Function and Carbon Storage**

3629 Impacts of extreme fires on biodiversity and rates of post-fire recovery have tended to require
3630 field-based approaches and are therefore not conducted routinely and at scale. Climate
3631 change, altered ignition, and suppression patterns are reshaping fire regimes and ecosystem
3632 functions, impacting biodiversity, ecosystem services, and C storage. Reduced seed quality,
3633 resprouting exhaustion, organic soil burning, and misaligned reproductive processes hinder
3634 post-fire regeneration (Burell et al., 2022; Nolan et al., 2021a; Johnstone et al., 2016). Other
3635 disturbances like drought and insect infestations compound these effects. For example,
3636 increased fire activity in boreal forests is reducing fire-adapted species like black spruce,
3637 favouring deciduous species and altering forest structure and C dynamics (Baltzer et al.,
3638 2021). In western North America, high-severity fires and warmer, drier conditions lead to poor
3639 forest recovery and transitions to shrublands or grasslands, affecting habitat, water regulation,
3640 and C storage (Coop et al., 2020). Similarly, in tropical savannas and forests, fire frequency
3641 can cause ecosystem transitions that are detrimental for C storage and biodiversity (Staver et
3642 al., 2011). Furthermore, release of ozone and particulate matter from fires including PM2.5
3643 negatively impacts plant and ecosystem health (Saxena et al., 2017), and reduces leaf area
3644 index through drought induced by aerosol radiative effects, leading to reduced carbon uptake
3645 (Tian et al., 2022).

3646 Observing and modelling shifting fire regimes face challenges due to variability, inconsistent
3647 historical data, and complex ecosystem responses. Short-term observations often miss long-
3648 term trends, and species-specific reactions are influenced by fire-adaptive traits and genetic
3649 variability (Grau-Andrés et al., 2024). Fire regimes respond to multiple disturbances,
3650 complicating analysis (Nolan et al., 2021a; Coop et al., 2020).

3651 Researchers are using new technologies to address these challenges. Advanced remote
3652 sensing, including satellite imagery and drones, provides detailed data on fire extent and
3653 severity (Burell et al., 2022; Baltzer et al., 2021). Improved modelling techniques, such as
3654 process-based simulation models and machine learning, enhance predictive capabilities
3655 (Coop et al., 2020; Nolan et al., 2021a). Long-term monitoring networks and citizen science
3656 initiatives contribute to comprehensive datasets (Baltzer et al., 2021; Coop et al., 2020).
3657 Advances in genomic tools, climate-adaptive management frameworks, and AI further
3658 improve fire impact predictions (Nolan et al., 2021a; Coop et al., 2020; Grau-Andrés et al.,
3659 2024). In all cases, efforts to harmonise local to regional-scale observations of impact are
3660 required, in a similar manner to emerging compilations of regional fire monitoring systems
3661 (e.g. Gincheva et al. 2024).

3662 A good example of emerging opportunities stemming from the assemblage of large datasets
3663 and application of AI is the comprehensive meta-analysis by Grau-Andrés et al. (2024), which
3664 analysed data from published studies to find that intensified fire regimes reduce plant
3665 abundance, diversity, and health globally. Increased fire severity has stronger effects than

3666 frequency, with forests, especially conifer and mixed forests, being more negatively impacted
3667 than open-canopy ecosystems. Woody plants are more susceptible than herbs due to slower
3668 growth and recovery rates. Arid and cold climates exacerbate these impacts, and transitions
3669 from surface to crown fires significantly reduce plant abundance, diversity, and health.

3670

3671 **8.2.6 Nature-based Solutions and Net Zero**

3672 The impacts of extreme fire on nature-based solutions such as reforestation projects are
3673 known to be of potential high importance, yet studies of these effects remain rare. Terrestrial
3674 ecosystems remove about 30% of annual anthropogenic C emissions through enhanced
3675 growth and ecosystem recovery, but land use change offsets this by increasing annual C
3676 emissions by about 12% (Friedlingstein et al., 2023). Nature-based solutions, including
3677 forestry projects for planting, restoration, or conservation, are prominent in net zero strategies
3678 (Seddon et al., 2020). However, these projects are at risk from fires, leading to inaccurate C
3679 accounting and reversal risks. The spatial clustering of C projects in specific areas can
3680 exacerbate risk due to climate variability, such as El Niño impacts. In C markets, projects often
3681 allocate a portion of their credits to a buffer pool to account for reversal risks, but this may not
3682 be sufficient in regions facing increasing wildfire extremes (Badgley et al., 2022; Anderegg et
3683 al., 2024).

3684 Despite these challenges, C markets offer opportunities to fund improved management of fire-
3685 prone ecosystems like savannas (Russel-Smith et al., 2015) and temperate forests (Nikolakis
3686 et al., 2022), benefiting ecosystems, climate, and local communities. Effective and scalable C
3687 markets require accurate and transparent monitoring systems for science-based fire
3688 management, precise C loss accounting from fires, and assessment of reversal risks. Novel
3689 satellite-based monitoring can provide early warnings, response to wildfire activity, and better
3690 estimates of ecosystem C stocks. Field studies remain essential for understanding the
3691 immediate and long-term impacts of fire on ecosystems (Silva et al., 2020), and new models
3692 forecasting fire risk over decades are needed to improve management strategies and support
3693 credible C claims.

3694

3695 **8.2.7 Water Quality and other Aquatic Impacts**

3696

3697 Impacts of extreme fires on water quality and other aquatic properties have tended to require
3698 field-based approaches and are therefore not conducted routinely and at scale. Fire impacts
3699 freshwater ecosystems mainly through (i) the loss of vegetation and litter cover and (ii) the
3700 enhanced input of soil, sediment, and ash. This leads to reduced rainfall interception,
3701 increased runoff, and greater rainfall reaching streams, lakes, and reservoirs (Smith et al.,
3702 2011). Increased runoff from burned hillslopes can cause erosion, debris flows, and localised
3703 flooding (Shakesby and Doerr, 2006).

3704

3705 Wildfire ash, enriched in nutrients and contaminants (e.g., metals, polycyclic aromatic
3706 hydrocarbons) compared to vegetation and soil (Bodi et al., 2014; Sánchez-García et al.,
3707 2023), affects water quality by increasing turbidity, temperature, nutrient, and toxin content,
3708 and decreasing dissolved oxygen. These changes can cause increased mortality in freshwater
3709 ecosystems, algal blooms, and water quality issues for supply catchments (Smith et al., 2011).
3710 For example, the 2016 Horse River Fire in Canada led to \$9 million in additional water
3711 treatment costs (Pomeroy et al., 2019). The impacts depend on the burned ecosystem type,
3712 fire size and severity, vegetation recovery rate, and post-fire rainfall patterns (Shakesby and
3713 Doerr, 2006; Nunes et al., 2018).

3714

3715 Fire emissions to the atmosphere contain compounds that can enrich or toxify aquatic
3716 ecosystems (Hamilton et al., 2022; Perron et al., 2022). While 2023 fire impacts on marine

3717 ecosystems are not yet well-documented, past extreme fires have disturbed ocean
3718 productivity far from the fire source. For instance, large Siberian fires in 2014 boosted Arctic
3719 phytoplankton blooms by adding reactive nitrogen (Ardyna et al., 2022). Similarly, 2019-2020
3720 Eurasian fires increased East Siberian Sea productivity by over 200% through nutrient
3721 deposition and ice melting (Seok et al., 2024). The 2017 Thomas fires in California and the
3722 2019/2020 Black Summer fires in Australia also caused significant changes in marine
3723 productivity and coastal ecosystems (Kramer et al., 2020; Ladd et al., 2023; Tang et al., 2021).
3724 The latter fueled a large phytoplankton bloom, temporarily offsetting the carbon emitted by the
3725 fires (Tang et al., 2021), though this effect fluctuates with each fire season (Hamilton et al.,
3726 2022; Wang et al., 2022).

3727

3728 **9 Competing Interests Statement**

3729

3730 SV is a member of the editorial board of Earth System Science Data. The authors declare no
3731 further conflict of interest.

3732

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3754

3755 **11 Author Contributions**

3756

3757 **Conceptualization:** Jones, M. W., Kelley, D. I., Burton, C. A., Di Giuseppe, F.

3758 **Project Administration:** Jones, M. W., Kelley, D. I., Burton, C. A., Di Giuseppe, F.

3759 **Data Curation:** Jones, M. W., Kelley, D. I., Burton, C. A., Di Giuseppe, F., Barbosa, M. L. F.,
3760 Brambleby, E., Lampe, S., Mataveli, G., McNorton, J., Spuler, F., Wessel, J., Parrington, M.

3761 **Formal Analysis:** Jones, M. W., Kelley, D. I., Burton, C. A., Di Giuseppe, F., Brambleby, E.,
3762 Lampe, S., Mataveli, G., McNorton, J., Spuler, F., Wessel, J., Parrington, M.

3763 Jones, M. W., Kelley, D. I., Burton, C. A., Di Giuseppe, F., Barbosa, M. L. F., Brambleby, E.,
3764 Hartley, A., Lampe, S., McNorton, J., Spuler, F., Wessel, J., Parrington, M., Hamilton, D. S.

3765 **Resources/Software:** Andela, N., Giglio, L., Parrington, M., van der Werf, G. R., Barbosa, M.
3766 L. F., Brambleby, E., Hartley, A., Lampe, S., McNorton, J., Spuler, F., Wessel, J., Burke, E.,
3767 San-Miguel-Ayanz, J., Qu, Y.

3768 **Visualisation:** Jones, M. W., Kelley, D. I., Burton, C. A., Di Giuseppe, F., Mataveli, G.,
3769 McNorton, J., Qu, Y., Lombardi, A.

3770 **Writing – Original Draft Preparation:** Jones, M. W., Kelley, D. I., Burton, C. A., Di Giuseppe,
3771 F., McNorton, J., Anderson, L., Archibald, S., Armenteras, D., Clarke, H., Doerr, S.,
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3773 V., Xanthopoulos, G., Carmenta, R., Hamilton, D. S., Hantson, S., Meier, S., Perron, M. M. G.,
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3775 **Writing – Review & Editing:** All authors.
3776

3777 **12 Data Availability**

3778

3779 BA data from NASA's MODIS BA product (MCD64A1) are extended from Giglio et al. (2018)
3780 and are available at Giglio et al. (2021, <https://lpdaac.usgs.gov/products/mcd64a1v061/>, last
3781 access: 9 July 2024). GFED4.1s fire C emissions data are extended from van der Werf and
3782 are available at <https://globalfiredata.org/> (last access: 9 July 2024). GFAS fire C emissions
3783 data are extended from Kaiser et al. (2012) and are available at
3784 [https://confluence.ecmwf.int/display/CKB/CAMS+global+biomass+burning+emissions+base
3785 d+on+fire+radiative+power+%28GFAS%29%3A+data+documentation](https://confluence.ecmwf.int/display/CKB/CAMS+global+biomass+burning+emissions+base+d+on+fire+radiative+power+%28GFAS%29%3A+data+documentation) (last access: 9 July
3786 2024). Global Fire Atlas are extended from Andela et al. (2019) and are available at Andela
3787 and Jones (2024, <https://doi.org/10.5281/zenodo.11400062>, last access: 9 July 2024).
3788 Regional summaries of the MODIS BA, GFED4.1s, GFAS, and the Global Fire Atlas are
3789 presented here are available at Jones et al. (2024, <https://doi.org/10.5281/zenodo.11400540>,
3790 last access: 9 July 2024). Studies utilising our regional summaries should cite both the current
3791 article and the primary reference for the variable(s) of interest: Giglio et al. (2018) for BA; van
3792 der Werf et al. (2017) for GFED4.1s fire C emissions; Kaiser et al. (2012) for GFAS fire C
3793 emissions; Andela et al. (2019) for the Global Fire Atlas.
3794

3795 Driving data, re-gridded BA target data for ConFire and ConFire outputs are available at Kelley
3796 et al. (2024, <https://doi.org/10.5281/zenodo.11420743>, last access: 9 July 2024). Historical
3797 (1960-2013) HadGEM3-A are available through the Centre for Environmental Data Analysis
3798 (CEDA) archive of the NERC's Environmental Data Service (EDS) at
3799 <http://catalogue.ceda.ac.uk/uuid/99b29b4bfeae470599fb96243e90cde3> (last access: 9 July
3800 2024). FireMIP / ISIMIP driving and output data is available from the Inter-Sectoral Impact
3801 Model Intercomparison Project (ISIMIP) repository at <https://data.ISIMIP.org/> (last access: 9
3802 July 2024).
3803

3804 **13 Code Availability**

3805

3806 ConFire attribution framework code (Kelley et al., 2021; Barbosa, 2024), was incorporated into
3807 the FLAME repository (https://github.com/douglask3/Bayesian_fire_models/tree/ConFire) and
3808 will be archived with a doi upon publication. Configuration settings for **Section 3** are in
3809 namelists/nrt.ini, while **Section 4** and **Section 5** are in namelists/ISIMIP.ini. Scripts for
3810 reproducing plots can be found in State of Wildfire GitHub repo (currently available at
3811 https://github.com/douglask3/State_of_Wildfires_report, with an archived doi available upon
3812 final publication).
3813

3814 The code used to produce the FWI attribution results is available in State of Wildfire GitHub
3815 repo (currently available at https://github.com/douglask3/State_of_Wildfires_report, with an
3816 archived doi available upon final publication). FWI code can be accessed via the ECMWF
3817 GitHub (<https://github.com/ecmwf-projects/geff>). Met Office implementation of the FWI is
3818 based on this code, and can be accessed at
3819 <https://github.com/MetOffice/impactstoolbox/tree/main> with registration. All details of the data
3820 and code used for the FireMIP attribution results is documented in Burton, Lampe et al. (2023).
3821

3822 The current version of ibicus, used for JULES-ES bias correction, is available from PyPI
3823 (<https://pypi.org/project/ibicus/>, last access: 9 July 2024) under the Apache License version

3824 2.0, and is described in detail in <https://ibicus.readthedocs.io/en/latest/> (last access: 9 July
3825 2024). The source code is available via GitHub (<https://github.com/ecmwf-projects/ibicus>, last
3826 access: 9 July 2024). Ibicus is also archived on Zenodo by Spuler and Wessel (2023;
3827 <https://doi.org/10.5281/zenodo.8101898>, last access: 9 July 2024). Model code and
3828 evaluation for bias-correction of JULES-ES model output can be found at the State of Wildfire
3829 GitHub repo (currently available at [https://github.com/jakobwes/State-of-Wildfires---Bias-](https://github.com/jakobwes/State-of-Wildfires---Bias-Adjustment)
3830 [Adjustment](https://github.com/jakobwes/State-of-Wildfires---Bias-Adjustment), with an archived doi available upon final publication).

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