



## 1 **Integration of global precipitation stable isotope data**

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10 **Abstract:** Precipitation plays a crucial role in the hydrological cycle and is vital for water resources  
11 management, climate change research, and ecosystem conservation. Precipitation stable isotopes  
12 serve as the "fingerprints" of precipitation, which can clearly trace the formation, transport, and  
13 subsequent processes of precipitation. However, due to the scarcity of precipitation stable isotope  
14 data, we face challenges of temporal discontinuity and spatial heterogeneity when studying it at  
15 large to medium scales. Therefore, we compiled precipitation hydrogen and oxygen stable isotope  
16 data ( $\delta^{18}\text{O}$  and  $\delta^2\text{H}$ ) from 2059 global sites spanning from 1961 to 2023, totaling 141,624 records.  
17 Our study indicates significant variations of global precipitation stable isotopes both spatially and  
18 temporally. Spatially, the isotopic composition of precipitation in different regions varies  
19 significantly due to factors such as geographical location, underlying surface conditions, and  
20 atmospheric circulation. Temporally,  $\delta^{18}\text{O}$  and  $\delta^2\text{H}$  show decreasing trends, while d-excess shows  
21 an increasing trend, with the impact of global temperature rise being very apparent. This  
22 precipitation stable isotope dataset provides robust support for our understanding of global  
23 precipitation changes and climate change. Through further investigation of precipitation stable  
24 isotope data, we hope to uncover more mechanisms and influencing factors of precipitation  
25 processes, providing a more accurate basis for the assessment and prediction of climate and water  
26 resource changes.

### 27 **1. introduction**

28 With the escalating global climate change, research on the impact of global water cycles and  
29 ecosystems on human society has become increasingly important (Anon, 2023; Intergovernmental



30 Panel on Climate Change (IPCC), 2023). Precipitation is one of the core components of the Earth's  
31 water cycle, and its spatial and temporal variations are crucial for accurate climate modeling and  
32 water resources management (Konapala et al., 2020; Bevacqua et al., 2022). Stable isotope  
33 technology has become an indispensable tool for understanding hydrological processes. Through  
34 analysis of the hydrogen and oxygen isotopic ratios in precipitation, key information such as the  
35 source, formation mechanism, and transport pathways of precipitation can be revealed (Cropper et  
36 al., 2021; Aggarwal et al., 2016). However, there are currently issues with the global precipitation  
37 stable isotope data, such as dispersion, inconsistency, and low accuracy, limiting our in-depth  
38 understanding of global climate and hydrological processes. Therefore, this study aims to integrate  
39 global precipitation stable isotope data, construct a more comprehensive dataset, fill gaps in existing  
40 research, and enhance our understanding of the Earth's hydrological cycle.

41 During the process of the hydrological cycle, stable isotope fractionation occurs with the phase  
42 change of water (Gat, 1996). Light isotopes ( $\delta^{16}\text{O}$  and  $\delta^1\text{H}$ ) are more prone to evaporate, while  
43 heavy isotopes ( $\delta^{18}\text{O}$  and  $\delta^2\text{H}$ ) are more easily enriched (Craig, 1961; Dansgaard, 1964). This  
44 difference leads to changes in the isotopic ratios in precipitation, which can be used to trace the  
45 source of water and hydrological processes. The stable isotope ratios in precipitation are influenced  
46 by temperature, known as the "temperature effect" (Juan et al., 2020). Generally, lower temperatures  
47 result in lower abundances of  $\delta^{18}\text{O}$  and  $\delta^2\text{H}$  in precipitation. The "temperature effect" manifests in  
48 two aspects: (1) evaporation from water vapor sources is reflected in the stable isotopes of  
49 precipitation in the area (Diekmann et al., 2021), (2) secondary evaporation affects precipitation  
50 condensation within clouds before reaching the ground, leading to enrichment of precipitation stable  
51 isotopes (more pronounced in arid regions) (Wang et al., 2016). Moreover, an increase in  
52 precipitation leads to a decrease in the stable isotope ratios in precipitation, as heavier isotopes are  
53 more easily removed during the precipitation process. This phenomenon is particularly evident  
54 during long-distance water vapor transport in the atmosphere. Additionally, an increase in  
55 precipitation weakens the "temperature effect" of precipitation stable isotopes. The variation in  
56 precipitation stable isotopic composition is also influenced by water vapor transport (Toride et al.,  
57 2021; Lekshmy et al., 2022). On one hand, there is a "continental effect" in the variation of  
58 precipitation stable isotopic composition (Winnick et al., 2014). On the other hand, the reverse  
59 altitude effect of water vapor stable isotopes is also reflected in the variation of precipitation stable



60 isotopes (Jing et al., 2022). This relationship makes stable isotopes a good proxy for climate. For  
61 example, precipitation stable isotopes are used to assess extreme weather events (Ansari et al., 2020;  
62 Lin et al., 2023).

63 The variations in precipitation isotopes can reveal multiple aspects of the water cycle, including  
64 the sources of precipitation, and the processes of evaporation and condensation. Thus, they are  
65 widely used in fields such as climate change research, hydrology, paleoclimatology, ecology, and  
66 geology (Zhang et al., 2021a; Caley et al., 2014). In climate change studies, stable isotopes of  
67 precipitation are used to track changes in past and present climate systems. By analyzing the isotopic  
68 ratios in ice cores, tree rings, and lake sediments (Jouzel et al., 1997), it is possible to reconstruct  
69 historical climate indicators such as temperature variations, precipitation levels, and evaporation  
70 intensity. This method provides an important tool for understanding the dynamics of the climate  
71 system, especially in interpreting past climate events in the context of global warming. The modern  
72 monitoring of precipitation stable isotopes began in 1961 with the observation network initiated by  
73 the International Atomic Energy Agency and the World Meteorological Organization (GNIP), which  
74 not only provides a powerful tool for understanding the current climate state but also validates  
75 reconstructions of past climates. Furthermore, isotope-supported atmospheric circulation models are  
76 becoming an important method for understanding modern and ancient climate changes (Bailey et  
77 al., 2018). The improvement of observational data is also clearly important for supporting the  
78 development of these models (Brady et al., 2019, p.1). In hydrology, stable isotope technology is  
79 used to study water cycle processes, such as the interactions between precipitation, surface water,  
80 and groundwater (Liberoff and Poca, 2023). Spatial and temporal variations in isotopic composition  
81 can help identify the sources and pathways of water, and assess the sustainability of water resources  
82 (Jiao et al., 2020). Through these isotopic studies, a deeper understanding of the dynamic changes  
83 in water resources and the complexity of the climate system can be achieved, providing a scientific  
84 basis for addressing environmental changes.

85 In this paper, we introduce the methods and data integration strategies employed in our study,  
86 including data sources, quality control, and statistical analysis. We present the constructed global  
87 dataset of stable isotopes in precipitation and provide preliminary analyses and validation of this  
88 dataset. Finally, we discuss the potential applications of the dataset and directions for future research,  
89 aiming to provide more reliable data support for global hydrological cycle studies and to offer



90 scientific foundations for addressing climate change and managing water resources.

## 91 **2. Data sources**

92 This manuscript compiles global precipitation stable isotope data from 1961 to 2023, which  
93 includes both field-collected and gathered data types. The collected data is sourced from the  
94 International Atomic Energy Agency's (IAEA) Global Network of Isotopes in Precipitation (GNIP)  
95 and the Water Isotope Website (<https://wateriso.utah.edu/waterisotopes/index.html>), encompassing  
96 a total of 141,624 data records (Fig. 1). The field-collected data originates from various locations  
97 across China, sampled by the Shiyang River Basin Observatory and analyzed by the Isotope  
98 Laboratory at Northwest Normal University, totaling 2,921 data records. Additionally, the  
99 meteorological data used for analysis in this paper is from TerraClimate—a gridded dataset of global  
100 terrestrial surface monthly climate and atmospheric water balance from 1958–2022 (Abatzoglou et  
101 al., 2018, p.1958–2015), available at: <https://www.climatologylab.org/terraclimate.html>. Global  
102 average temperature data is provided by NASA's Goddard Institute for Space Studies (GISS).  
103 Marine meteorological data from NCEP-NCAR Reanalysis 1 and NOAA Extended Reconstructed  
104 Sea Surface Temperature (SST) V4.

## 105 **3. Data processing steps and quality control**

106 Data Collection: The collected data contained many missing values, anomalies, duplicate  
107 entries, as well as missing dates, and errors or absences in geographic coordinates (latitude and  
108 longitude). Consequently, the raw data was preprocessed and cleaned. Missing data was interpolated,  
109 and data entries that could not be completed, as well as duplicate data, were deleted.

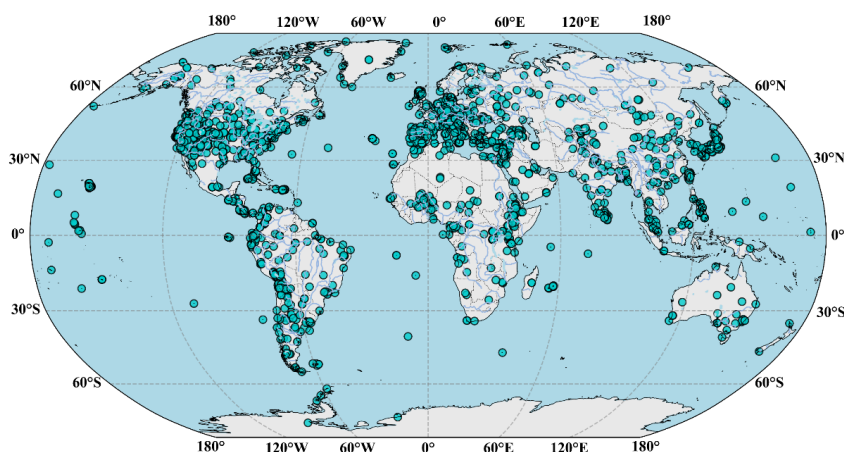
110 Measured Data: Precipitation samples were collected using a standard rain gauge. Immediately  
111 after each precipitation event, the collected precipitation samples were transferred to 100ml high-  
112 density sample bottles. To prevent data errors due to evaporation, the collected water samples were  
113 stored in a refrigerator at approximately 4°C. At the start of the experiments, the precipitation  
114 samples were naturally thawed at room temperature. During sampling, a 0.45µm filter was used to  
115 filter impurities, and the samples were transferred to 2ml sample bottles. The isotopic values were  
116 determined using a Liquid Water Isotope Analyzer (DLT-100, Los Gatos Research, USA). For each  
117 sample, six measurements were taken, discarding the first two results to avoid cross-sample  
118 influence. For values that were anomalies or not verified by the LWIA post Analysis software, we  
119 reselected parallel samples for measurement to ensure the accuracy of the data. The isotopic



120 abundances of  $^{18}\text{O}$  and  $^2\text{H}$  are denoted using the delta ( $\delta$ ) symbol relative to the Vienna Standard  
121 Mean Ocean Water (V-SMOW) standard, following the equation:

122 
$$\delta_{\text{sample}} (\text{‰}) = \left[ \frac{R_{\text{sample}}}{R_{V\text{-smow}}} - 1 \right] \times 1000$$

123 where ( R ) is the ratio of the heavier isotope to the lighter isotope (i.e.,  $^{18}\text{O}/^{16}\text{O}$ ) or ( $^2\text{H}/^1\text{H}$ ).  
124 We have validated our isotopic measurements against the International Atomic Energy Agency  
125 (IAEA) standard (V-SMOW2) to ensure the comparability of isotopic measurements across different  
126 laboratories and instruments.



127  
128 Fig. 1 Global distribution of precipitation stable isotope sites

## 129 4. Results and discussion

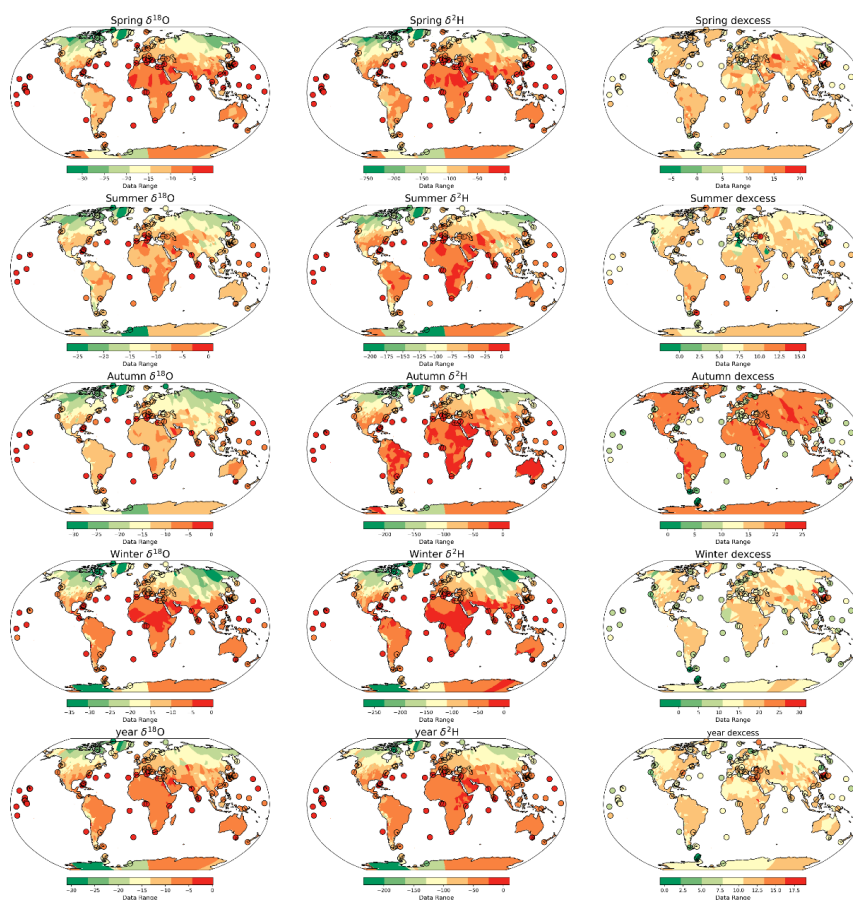
### 130 4.1 Spatial distribution of stable isotopes in precipitation

131 The spatial distribution of global terrestrial  $\delta^2\text{H}$  and  $\delta^{18}\text{O}$  exhibits a clear latitudinal correlation,  
132 characterized by isotopic values becoming progressively depleted from low to high latitudes (Fig.  
133 2). This variation is primarily controlled by temperature, especially on a global scale. In the low  
134 latitudes, often influenced by tropical rainforest and monsoon climates, the high temperatures and  
135 humidity intensify the "amount effect" in precipitation's stable isotopes, while weakening the  
136 "temperature effect" (Eastoe and Dettman, 2016). Conversely, in the arid climates of northern Africa,  
137 the Arabian Peninsula, and Central Asia, the high temperatures and dry conditions promote  
138 evaporation fractionation and "below-cloud secondary evaporation," leading to relatively enriched  
139 isotopic values in precipitation in these regions. In mid to low latitude mountainous areas, such as



140 the Tibetan Plateau, the Alps, and the Andes in South America, precipitation stable isotope ratios  
141 are generally low, reflecting the unique climatic and precipitation processes at high altitudes.  
142 Additionally, in Western Europe, particularly in coastal areas near the North Atlantic, the isotopic  
143 high values extend to higher latitudes, reflecting the region's higher temperatures. Seasonally, the  
144 Northern Hemisphere experiences lower stable isotope values in precipitation during winter and  
145 spring, while values are relatively enriched during summer and autumn, correlating with seasonal  
146 climate changes. The Southern Hemisphere presents a unique case, as most regions are tropical and  
147 land areas are relatively smaller, leading to less pronounced seasonal variations in precipitation  
148 stable isotopes. The composition of stable isotopes in global precipitation is influenced by multiple  
149 factors, including latitude, climate type, seasonal variations, and topography, which together  
150 determine its complex spatial and temporal distribution patterns.

151 In the oceans near the equator, we observe high values of  $\delta^2\text{H}$  and  $\delta^{18}\text{O}$  isotopes, which directly  
152 reflect the region's high evaporation rates and rapid moisture exchange. In contrast, in polar regions,  
153  $\delta^2\text{H}$  and  $\delta^{18}\text{O}$  are generally lower (Fig. 2). This distribution pattern is consistent with the spatial  
154 differentiation of stable isotopes' "latitudinal gradient" on land. Specifically, the North Atlantic and  
155 Western Pacific regions exhibit isotopic enrichment or depletion distinct from other areas at similar  
156 latitudes, revealing possible climatic peculiarities or special atmospheric circulation patterns in  
157 these regions. In the North Atlantic, the observed isotope values are higher than in other areas at the  
158 same latitude, which can be attributed to the intensified non-stationary climate variations due to the  
159 North Atlantic Oscillation (NAO) (Schurer et al., 2023). Additionally, it is associated with rising sea  
160 surface temperatures (Karnauskas et al., 2021). In contrast, in Southeast Asia, observed isotope  
161 values are lower than in other regions at the same latitude, which correlates with increased  
162 precipitation in the area. With global warming, we not only observe an increase in total precipitation  
163 but also an intensification of precipitation extremes (Zhang et al., 2021b). Due to the tropical  
164 monsoon climate and strong convective weather systems, the region experiences more rainfall,  
165 which in turn enhances the "amount effect" of stable precipitation isotopes.



166

167 Fig. 2 Spatial distribution of precipitation stable isotopes (where dots represent isotope values at  
168 sampling sites near the ocean and coast.)

#### 169 4.2 Stable isotope time series of precipitation in different climatic zones

170 Based on the Köppen climate classification (Beck et al., 2018), time series of the stable isotopic  
171 values of precipitation (weighted averages) were constructed for different climatic zones (Fig. 3).  
172 Tropical (A) and desert climates (B) exhibit high values of stable isotopes in precipitation. Since  
173 1960, the tropical and temperate climate zones have shown steady fluctuations in stable isotopic  
174 values of precipitation, while arid climate zones have experienced significant volatility. Without  
175 considering the changes in stable isotopes under extreme conditions, this suggests that tropical and  
176 temperate climates have a relatively stable hydrothermal combination relationship. During the  
177 period 1985-1995, the stable isotopic values of desert climate precipitation were lower than those  
178 of the temperate zone, indicating a trend of decreased temperatures or increased precipitation in



179 desert climates during this time. Polar (D) and arctic climates (E) exhibit lower stable isotopic values  
180 in precipitation compared to other climatic zones. There is a process of hydrothermal exchange  
181 between high and low latitudes; the high temperatures at low latitudes cause the precipitation  
182 reaching the ground to be enriched in heavy isotopes, and as water vapor is transported to higher  
183 latitudes, the heavy isotopes are progressively stripped away. Additionally, the lower temperatures  
184 in high-latitude areas result in weaker sub-cloud secondary evaporation. Therefore,  $\delta^{18}\text{O}$  and  $\delta^2\text{H}$   
185 in high-latitude precipitation are relatively depleted. From 1975 to 1985, cold climate zones showed  
186 a continuous increase in  $\delta^{18}\text{O}$  and  $\delta^2\text{H}$  values, and an upward trend has also been observed since  
187 2015, indicating an intensifying influence of temperature in cold and arctic climate zones. It can be  
188 inferred that under the backdrop of global warming, cold and arctic climate zones might contribute  
189 more water vapor to the global hydrological cycle, significantly impacting global climate and water  
190 resource patterns. These analyses may obscure some information (e.g., the classification of climate  
191 zones also varies by distance, such as between the Northern and Southern Hemispheres), leading to  
192 less apparent linear trends in isotopic changes across different climates. However, the dataset of  
193 stable isotopes in precipitation we constructed also reflects the differences among various climatic  
194 zones, allowing for more precise tracing of global precipitation and a deeper understanding of the  
195 global water cycle.

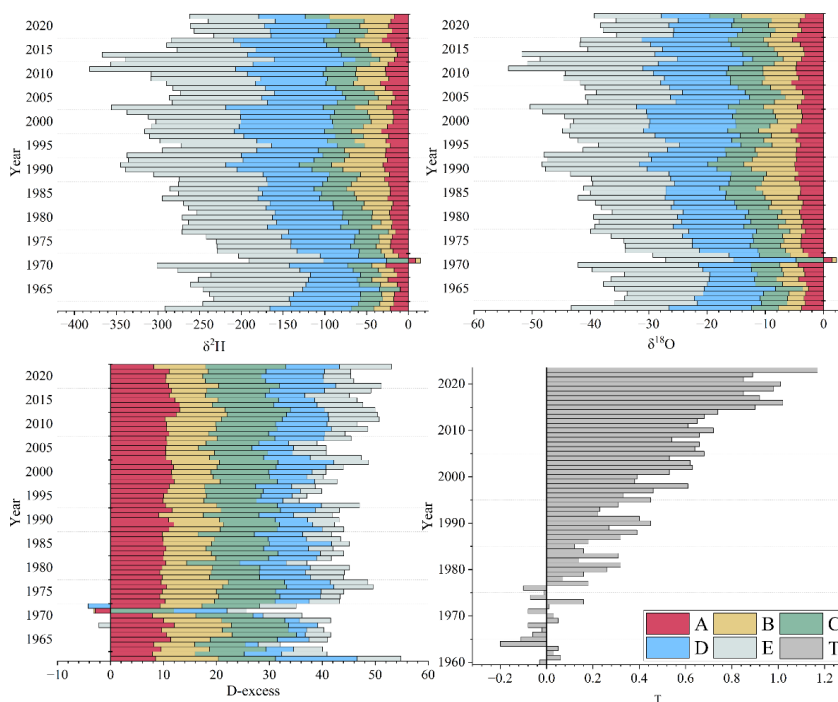
196 The ocean is the core source of global terrestrial precipitation, and over the past several decades,  
197 the isotopic ratios of  $\delta^2\text{H}$  and  $\delta^{18}\text{O}$  over the ocean have exhibited significant interannual fluctuations,  
198 showing a general declining trend (Fig. 4). At the same time, the d-excess values have displayed an  
199 increasing trend. A possible explanation for this phenomenon is that regional differences in  
200 precipitation rates lead to a decrease in the stable isotopic ratios of precipitation, while the dominant  
201 trend remains driven by isotopic fractionation due to rising temperatures.

202 Stable isotopes in precipitation are positively correlated with atmospheric temperature changes  
203 recorded in historical climate data. This correlation demonstrates the potential of stable isotopes as  
204 indicators of global climate change, particularly in identifying the response of precipitation  
205 processes and water vapor sources to climatic factors. However, the years with rising temperatures  
206 do not always correspond to the years with changes in stable isotopic ratios (weak correlation),  
207 reflecting the complexity of factors affecting stable isotopes in precipitation, indicating that a single  
208 indicator (such as global average temperature) cannot fully explain climate changes in all regions.

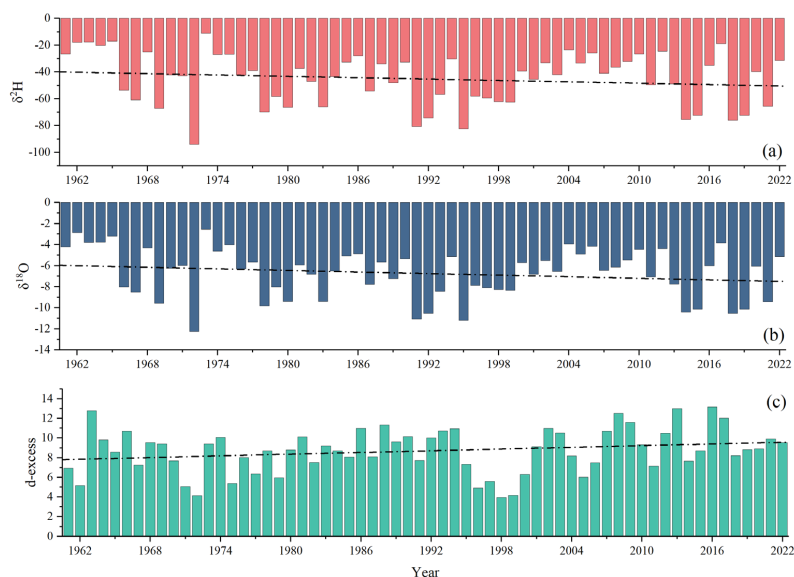




209 Stable isotopes can serve as comprehensive climate proxy indicators. Moreover, d-excess exhibits  
210 temporal characteristics different from those of  $\delta^2\text{H}$  and  $\delta^{18}\text{O}$ , reflecting changes in different  
211 atmospheric water vapor sources and dry-wet conditions.



212  
213 Fig. 3 Temporal trends of stable isotopes of precipitation in different climatic zones, A for tropical  
214 climate, B for arid climate, C for temperate climate, D for boreal climate, E for polar climate, and  
215 T for global average temperature ( $^{\circ}\text{C}$ ).



216

217 Fig. 4 Temporal trends in stable isotopes of precipitation over the global ocean

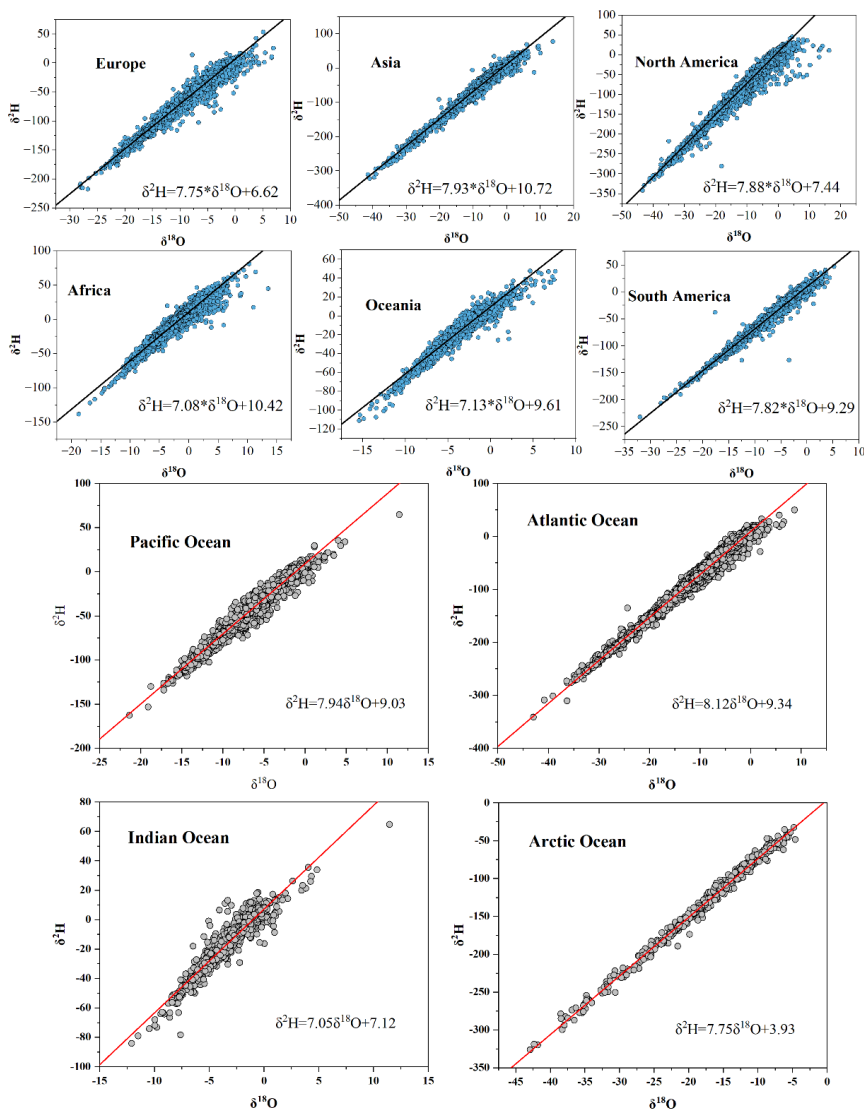
### 218 4.3 Local Meteoric Water Lines in different regions

219 The Global Meteoric Water Line (GMWL) describes the global annual average relationship  
220 between the ratios of hydrogen and oxygen isotopes ( $\delta^{18}\text{O}$  and  $\delta^2\text{H}$ ) in precipitation. GMWL was  
221 originally defined by Harmon Craig in 1961 as  $\delta^2\text{H} = 8\delta^{18}\text{O} + 10$ . GMWL is considered the  
222 "expected" equilibrium relationship, assuming the slope is generated by the ratio of equilibrium  
223 fractionation factors. Based on this dataset, the fitted equation for the global atmospheric  
224 precipitation line is:  $\delta^2\text{H} = 7.95\delta^{18}\text{O} + 9.40$ . The lower intercept and slope compared to the past  
225 might indicate a change in the temperature environment of evaporation and condensation in the  
226 precipitation process, which could be related to changes in the global average temperature (Fig. 3d).  
227 Factors such as accelerated hydrological cycles, climate models, and changes in precipitation  
228 sources could have impacts (Dangendorf et al., 2019; Koutsoyiannis, 2020). On one hand, a faster  
229 conversion from evaporation to precipitation and shorter atmospheric residence time might reduce  
230 the polarization of fractionation effects, partly explaining the decreased slope. On the other hand,  
231 evaporation from oceans generally contains more heavy isotopes, altering the global average isotope  
232 ratio.

233 The slope and intercept variability of Local Meteoric Water Lines (LMWL) across different  
234 continents and oceans (excluding Antarctica) reflect the diversity of hydroclimatic conditions and



235 isotopic fractionation (Putman et al., 2019) (Fig.5). Fitted using the least squares method, these  
236 precipitation lines' variability in slope and intercept are primarily driven by humidity changes and  
237 also reflect the combined effects of temperature, precipitation, and other meteorological factors,  
238 revealing regional climatic differences. In Asia, the slope of LMWL is closest to the global average  
239 (8‰), indicating a more uniform hydrothermal combination under its temperate climate conditions.  
240 In contrast, Africa's LMWL slope is the lowest, closely related to its extensive arid and tropical  
241 climates. The stable isotopes of oceanic precipitation exhibit characteristics of equilibrium  
242 fractionation, particularly in the Pacific and Atlantic Oceans, which not only span multiple thermal  
243 zones but also serve as the primary moisture sources for most terrestrial regions. The Indian Ocean's  
244 LMWL slope is lower due to most sampling points being near the equator, mainly related to its low-  
245 latitude temperature characteristics. The slopes and intercepts inside the polar circles are lower, a  
246 phenomenon associated with the global warming trend, which promotes the anomalous performance  
247 of stable isotopes in high-latitude precipitation, further substantiating our earlier inference about  
248 increased moisture supply from cold and polar regions under global warming.



249

250 Fig. 5 Local atmospheric precipitation lines on land and at sea

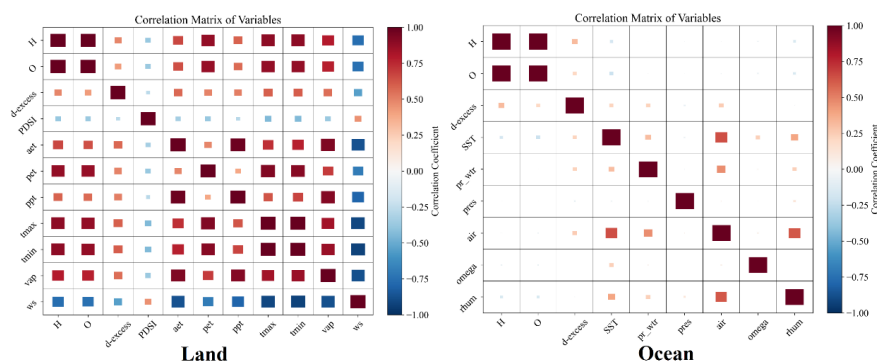
#### 251 4.4 Precipitation stable isotopes and meteorological elements

252 By analyzing the correlations between stable precipitation isotope variables ( $\delta^{18}\text{O}$ ,  $\delta^2\text{H}$ , D-  
253 excess) and key meteorological elements (including wind speed, vapor pressure, highest and lowest  
254 temperatures, precipitation, potential and actual evapotranspiration) (Fig. 6), results indicate  
255 significant negative correlations between the isotopic variables  $\delta^{18}\text{O}$  and  $\delta^2\text{H}$  with wind speed (ws)



256 and vapor pressure (vap), with correlation coefficients of -0.76 and -0.78, respectively. This suggests  
257 that under conditions of increasing wind speeds and vapor pressures, the proportions of hydrogen  
258 and oxygen isotopes in precipitation are significantly reduced. This reflects the enhanced  
259 fractionation effects of isotopes during increased evaporation and the mixing process of atmospheric  
260 moisture. Particularly, the negative correlation of D-excess with wind speed (-0.48) is lower than  
261 that of  $\delta^{18}\text{O}$  and  $\delta^2\text{H}$ , pointing to D-excess being controlled by multiple factors, including the initial  
262 isotopic composition of water vapor and its subsequent transformation in the atmosphere. In contrast,  
263 the correlation of D-excess with meteorological variables is generally weaker, reflecting the  
264 uniqueness and complexity of D-excess in the evaporation and precipitation recycling process.  
265 Additionally, precipitation (ppt) shows a positive correlation with potential and actual  
266 evapotranspiration (pet and act), suggesting that increased precipitation may lead to higher water  
267 availability, thus enhancing the surface's capacity for evapotranspiration. This also demonstrates  
268 that stable precipitation isotopes are highly reliable indicators for assessing global  
269 evapotranspiration. Moreover, the relative independence of D-excess from meteorological elements  
270 offers a unique perspective for exploring different mechanisms in water vapor sourcing and the  
271 water cycle process. These findings provide valuable insights for further understanding the  
272 application of isotope data in environmental and climate research.

273 Over the oceans, the correlation between precipitation stable isotopes and meteorological  
274 variables is generally weak due to several factors. Firstly, global weather systems dictate the  
275 distribution of precipitation stable isotopes. Secondly, extreme weather events over the oceans, such  
276 as typhoons and intense convective events (Wang et al., 2021), significantly impact the fractionation  
277 of stable isotopes, resulting in notable variations in isotopic compositions under these conditions.  
278 These factors collectively contribute to greater spatial variations in precipitation stable isotopes over  
279 the oceans compared to those on land, with the influence of specific regional meteorological factors  
280 being less apparent. Moreover, compared to terrestrial areas, the scarcity of data over the ocean  
281 limits a detailed understanding and precise matching of the relationship between isotopes and  
282 meteorological elements, thus affecting the assessment of their correlation. The ocean is a crucial  
283 component of the global water cycle, and enhancing the observation of stable isotopes in oceanic  
284 precipitation can not only provide more detailed clues about global climate change but also help  
285 improve the accuracy of assessments of global climate trends and water resource distribution.



286

287 Fig. 6 Correlation of precipitation stable isotopes with meteorological elements

288 **5. Summary and outlook**

289 The global dataset of precipitation stable isotopes is a highly valuable tool for analyzing and  
290 simulating trends in climate and water resources. By relying on this re-integrated dataset, we can  
291 detect the impacts of climate change on precipitation quantity, frequency, and spatial distribution,  
292 thereby gaining a deeper understanding of the evolutionary trends in the global climate system.  
293 Precipitation stable isotope data is not only useful for monitoring climate change but also provides  
294 critical support for studies on the hydrological cycle. Through dataset analysis, we can understand  
295 the sources, transport pathways, and deposition processes of precipitation within the hydrological  
296 cycle, revealing the patterns and influencing factors of hydrological cycles in different regions. This  
297 new dataset is significant for water resource management. By analyzing precipitation stable isotope  
298 data, we can assess the availability of water resources, guide their rational development, utilization,  
299 and conservation. Additionally, this dataset can be used to optimize water resource allocation  
300 schemes to address the challenges brought about by climate change and human activities on the  
301 balance of water supply and demand. We hope to further utilize these data resources to deepen our  
302 understanding of the complexity of global climate and hydrological cycles, providing scientific  
303 decision support for addressing climate change and achieving sustainable water resource  
304 management. Moreover, compared to terrestrial areas, there is relatively less precipitation stable  
305 isotope data over the oceans, and future efforts should focus on enhancing the observation of stable  
306 isotopes in oceanic precipitation.

307 **Data Availability**

308 Data described in this manuscript can be accessed at Mendeley Data under



309 10.17632/9gxtc4xzv9.1 (Zhu, 2024).

#### 310 **Author contribution Statement**

311 Longhu Chen: Conceptualization and writing-original draft preparation; Qinqin Wang: Data  
312 processing; Guofeng Zhu: Writing review and editing; Rui Li: Methodology; Siyu Lu: Experiment;  
313 Xinrui Lin: Validation; Dongdong Qiu: Methodology; Gaojia Meng: Visualization; Yinying Jiao:  
314 Data processing; Yuhao Wang: Visualization; Jing Liu: Modification; Yutong He: Modification;  
315 Yanan Li: Modification.

#### 316 **Declaration of Interest Statement**

317 The authors declare that they have no known competing financial interests or personal  
318 relationships that could have appeared to influence the work reported in this paper.

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