

1 **High-resolution Carbon cycling data from 2019 to 2021 measured at six**  
2 **Austrian Long-Term Ecosystem Research sites**

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18

19 Abstract

20 Seven long-term observation sites have been established in six regions across Austria, covering major  
21 ecosystem types such as forests, grasslands and wetlands across a wide bioclimatic range. The  
22 purpose of these observations is to measure key ecosystem parameters serving as baselines for  
23 assessing the impacts of extreme climate events on the carbon cycle. The data sets collected include  
24 meteorological variables, soil microclimate, CO<sub>2</sub> fluxes and tree stem growth, all recorded at high  
25 temporal resolution (15 – 60 minutes) between 2019 and 2021 (including one year of average  
26 climate conditions and two comparatively dry years). The DOIs of the dataset can be found in the  
27 data availability chapter. The sites will be integrated into the European Research Infrastructure for  
28 Integrated European Long-Term Ecosystem, Critical Zone, and Socio-Ecological Research (eLTER RI).  
29 Subsequently, new data covering the variables presented here will be continuously available through  
30 its data integration portal. This step will allow the data to reach its full potential for research on  
31 drought-related ecosystem carbon cycling.

32

33 1. Introduction

34 Climate change has been affecting ecosystems globally with strong implications for the terrestrial  
35 carbon (C) cycle, which in turn feeds back to the climate system (Heimann and Reichstein, 2008). As  
36 an emerging feature of climate change, extreme climatic events (ECEs) are expected to occur with  
37 increasing frequency and intensity in the coming decades (IPCC, 2021). ECEs are considered to exert  
38 stronger impacts on ecosystems and the services they provide to mankind than gradual changes in  
39 climate (Frank et al., 2015; Reichstein et al., 2013; Grünzweig et al., 2022; Anderegg et al., 2020).  
40 Understanding, predicting and managing extreme climate events and their consequences for

41 ecosystems and societies will therefore be one of the big challenges in the coming decades. To detect  
42 and attribute impacts of ECEs on ecosystem processes and services they need to be evaluated on the  
43 background of the typical interannual range of these processes (Ciais et al., 2005; Bernal et al., 2012;  
44 Fu et al., 2020; Schindlbacher et al., 2012) and analyses of ecosystem resilience to ECEs require a  
45 robust quantification of baselines of ecosystem functioning (Bahn and Ingrisch, 2018; Ingrisch and  
46 Bahn, 2018). For deriving such baselines as well as interannual variability of ecosystem carbon  
47 cycling, coordinated and representative observation networks need to be in place to enable data  
48 retrieval as well as rapid-response scientific campaigns to study after-effects and post-disturbance  
49 trajectories resulting from ECEs (Kulmala, 2018; Mahecha et al., 2017; Mirtl et al., 2018; Dirnböck et  
50 al., 2019; Müller and Bahn, 2022). Datasets obtained through such observation networks are also  
51 essential for benchmarking models (Futter et al., 2023; Baatz et al., 2021; Wu et al., 2018) and for  
52 comparison with ecosystem experiments (Kröel-Dulay et al., 2022).

53 Within a research infrastructure project focusing on ecosystem carbon, nitrogen, and water fluxes  
54 (Long-Term Ecosystem Research for Carbon, Water, and Nitrogen (LTER-CWN, [https://www.lter-  
55 austria.at/cwn/](https://www.lter-austria.at/cwn/)), we equipped seven long-term observation sites in six regions, which are part of the  
56 existing Long-Term Ecological Research Network of Austria (LTER-Austria), with high temporal  
57 resolution (30-60 minutes) C cycle measurements. The sites cover three major ecosystem types  
58 occurring across Austria (forests, managed mountain grassland, wetlands) and most of them are part  
59 of socio-ecological research platforms for transdisciplinary studies (Figure 1). Here, we  
60 provide observational ecosystem response data capturing naturally-occurring ECEs from the first  
61 three years after the onset of the infrastructure, 2019 to 2021. These data sets include  
62 meteorological variables, soil microclimate, CO<sub>2</sub> flux measurements using automated chambers (soil  
63 CO<sub>2</sub> efflux) and eddy covariance techniques (net ecosystem exchange), respectively, and tree stem  
64 radial increments and shrinkage in forested plots.

65

## 66 2. Site descriptions

67 The sites are key research infrastructures for ecosystem-related greenhouse gas observations in  
68 Austria. They include forests (Klausen-Leopoldsdorf and Rosalia in Lower Austria, Zöbelboden in  
69 Upper Austria, and Stubai in Tyrol), mountain grassland (Stubai, Tyrol), and wetlands (Pürgschachen  
70 Moor, Styria and Lake Neusiedl reed belt, Burgenland). This network of sites covers typical forest,  
71 alpine and wetland ecosystems of Central Europe (Figure 1). Furthermore, the sites represent  
72 different geological characteristics, from crystalline rock in the central Alps to the limestone in the  
73 northern Alps to unconsolidated Holocene sediments in lowlands. All sites are part of the Austrian  
74 LTER network and, once officially launched, will be included in the European eLTER research  
75 infrastructure (<https://elter-ri.eu/>). For a detailed description of the sites, we refer to the Dynamic  
76 Ecological Information Management System - Site and dataset registry (DEIMS-SDR) (Table 1).

### 77 2.1. Rosalia Forest Demonstration Centre (Mixed beech forest)

78 The Rosalia Forest Demonstration Centre was settled in 1972, as a cooperation between the BOKU  
79 University, Vienna, Austria, and the Austrian Federal Forests, and has approximately 1000 ha in the  
80 western slopes of the Rosalia Mountains (Rosaliengebirge) in Lower Austria (Figure 1, Table  
81 1). The forest hosts all major tree species occurring in Austria, i.e. European beech (*Fagus sylvatica*  
82 L.), Norway spruce (*Picea abies* (L.) H.Karst.), Scots pine (*Pinus sylvestris* L.), Larch (*Larix decidua*  
83 Mill.), and Fir (*Abies alba* Mill.). The altitude ranges from 320 to 725 m a.s.l., and mean annual  
84 temperature and mean annual precipitation are 6.5 °C and 796 mm, respectively. Substrate is mainly  
85 composed by crystalline rocks, and soils are predominantly cambisols (Working Group WRB 2015);

86 sporadically in combination with planosols (in plains and moderate slopes), with fluvisols (in valleys)  
87 or podzolic cambisols (steep slopes) (Fürst et al., 2021).

88 The demonstration forest holds several experimental and observation sites distributed along its area,  
89 including water, soil, vegetation and air observations (e.g. Gillespie et al., 2023). A watershed (220  
90 ha) is subject to hydrological observations (Fürst et al., 2021), and the forest is regularly monitored  
91 on permanent plots (Gollob et al., 2020). The meteorological data presented here originates from  
92 three stations located at 385 (Mehlbeerleiten), 500 (Kuhwald) and 640 m a.s.l. (Heuberg). The C cycle  
93 data was measured on a long-term experimental site launched in 2012. The site is located in a pure  
94 mature beech stand at 600 m a.s.l. (47° 42' 26" N; 16° 17' 59" E). It faces north-west, with a slope of  
95 approximately 20 %. This experiment focuses on investigating the effect of changing precipitation  
96 patterns ~~for selected on soil~~ nitrogen fluxes, soil microbial changes, greenhouse gas efflux, and soil  
97 water processes biogeochemical and microbiological processes (Leitner et al., 2017; Liu et al., 2019;  
98 Schwen et al., 2015; Gillespie et al., 2024). Monitoring is performed on control and on manipulated  
99 plots. The data from both natural and manipulated plots is published with this paper. Manipulation  
100 involves the use of rain-out-shelters (for simulating drought periods of different length) and of an  
101 irrigation system (for recreating rainfall events of different intensity). The monitoring infrastructure  
102 involves the measurements of greenhouse gases (GHG) (N<sub>2</sub>O, CH<sub>4</sub> and CO<sub>2</sub>) fluxes, soil nutrients  
103 (suction cups) and microclimate parameters.

#### 104 2.2. Klausen-Leopoldsdorf (Beech forest)

105 The site, Klausen-Leopoldsdorf, is located about 40 km south-west of Vienna on a NNE-facing slope  
106 and was founded in the 1990s as one of Austria's site contributing to the International Co-operative  
107 Programme on Assessment and Monitoring of Air Pollution Effects on Forests (ICP Forests) (Neumann  
108 and Starlinger, 2001). The site is divided into four different sub-areas within a small catchment: 1)  
109 the ICP Forests site, 2) a weather station, located 2.7 km from the ICP Forests intensive plots at 398  
110 m a.s.l., 3) a catchment runoff weir (475 m a.s.l.), and 4) the LTER-CWN measurement plot (520 m  
111 a.s.l.), where the C-cycle data presented here was measured (~~Figure 1~~Figure 1, Table 1). The forest  
112 within the measurement plot is a pure beech (*Fagus sylvatica* L.) stand. The mean annual  
113 temperature is 8°C, mean annual precipitation is 801 mm (2010-2022). The geological substrate is  
114 sandstone, the soil type is mainly stagnic cambisol/dystric cambisol (Working Group WRB 2015).  
115 Instruments installed on the LTER-CWN measurement area include a sap flow and dendrometer  
116 measurement system on 10 trees, 12 GHG automated measurement chambers for CO<sub>2</sub> respiration,  
117 soil moisture and soil temperature sensors in different soil depths (5 – 30 cm).

118 In addition to the data presented here, many other data sets are available. Soil GHG fluxes (manual  
119 sampling) were measured starting in the year 2001 (Kitzler et al., 2006). On the ICP Forests site,  
120 instruments for long-term monitoring (since 1996) such as soil moisture, air temperature and relative  
121 humidity, soil temperature, soil solution with suction cups, throughfall deposition, litterfall traps,  
122 stemflow, and manual and automatic dendrometers are installed and the data is available under  
123 <https://bfw.ac.at/lims/level2.daten> or via the ICP Forests Program Centre.

#### 124 2.3. Lake Neusiedl (reed belt)

125 The measurement site is located in the eastern reed belt of the lake and as such inside the National  
126 Park Lake Neusiedl - Seewinkel (~~Figure 1~~Figure 1, Table 1). The region (average altitude: 120 m.a.s.l.)  
127 is characterized by a (sub)-continental Pannonian climate with a mean annual precipitation of 576  
128 mm (2013-2022). The reed belt is a dynamic ecosystem consisting of a mosaic of reed stocks  
129 (*Phragmites australis* (Cav.) Trin. ex Steud.), sediment and open water areas. Increasing dry periods  
130 and thus successive drying of the reed belt since 2018 have led to an increase in reed stocks within

131 the belt, as well as an increase in sediment areas and a strong decline in open water areas, according  
132 to a 2021 study that investigated the spatial and temporal variations within the reed ecosystem at  
133 Lake Neusiedl (Buchsteiner et al., 2023). Processes driving CH<sub>4</sub> emissions from the reed belt have  
134 recently been investigated in detail (Baur et al., 2024).

135 The data presented here stems from devices permanently installed on site. They include an eddy  
136 covariance tower for CO<sub>2</sub>, CH<sub>4</sub> and water vapor fluxes and relevant accompanying meteorological  
137 parameters as well as soil heat flux, soil moisture, and soil temperature sensors.

#### 138 2.4. Pürgschachen Moor (peat bog)

139 The Pürgschachen Moor is located on the bottom of the Styrian Enns valley at an altitude of 632 m  
140 a.s.l. (Figure 1, Table 1). It is a pine peat bog with an extent of about 62 ha. Thus, it is the  
141 largest (to a large part) intact valley peat bog in Austria with a closed peat moss cover and a good  
142 example of the formerly widely distributed peatlands of inner-alpine valleys of the European Alps.  
143 The mean average temperature is 8.2 °C and mean annual precipitation is 1233 mm (2013-2022). The  
144 typical vegetation of the peat bog is constituted of three associations of plants *Pino mugo-*  
145 *Sphagnetum magellanici* (pine peat bog association), *Sphagnetum magellanici* (coloured bog moss  
146 association), and *Caricetum limosae* (bog sedge association), depending on the prevailing  
147 hydrological site conditions. The current mean water table depth is about 14 cm below soil surface at  
148 the central peat bog area. Peat decomposition and related CO<sub>2</sub> and CH<sub>4</sub> fluxes were subject of a  
149 series of research studies (Drollinger et al., 2019; Knierzinger et al., 2020; Müller et al., 2022; Glatzel  
150 et al., 2023).

151 The data presented here stems from devices permanently installed roughly in the center of the peat  
152 bog. They include an eddy covariance tower for CO<sub>2</sub>, CH<sub>4</sub> and water vapor fluxes and relevant  
153 accompanying meteorological parameters as well as soil heat flux, soil moisture, and soil  
154 temperature sensors.

#### 155 2.5. Stubai (subalpine hay meadow, Larch and Spruce forest)

156 The two observation plots used in this study are part of the LTER Site Stubai (Table 1), which is  
157 located in the Stubai Alps in Tyrol, Austria (Figure 1). Research at the study site was  
158 established in 1993. The two observation plots are a mountain grassland and a subalpine forest at an  
159 alpine pasture area called "Kaserstattalm". The underlying rock is siliceous and calcareous. The  
160 average air temperature is about 3°C and the precipitation approx. 1100 mm. About 35% of the  
161 annual precipitation occurs as snow during winter months.

162 The grassland site is located at an altitude of 1810 -1850 m a.s.l on a south-east facing slope with an  
163 inclination of ca. 20°. The site is an extensively managed meadow that is harvested once a year in  
164 early August and grazed lightly in late summer. The soil is a dystric cambisol (Working Group WRB  
165 2015). The vegetation type is a *Trisetetum flavescens* and consists of perennials grasses and forbs  
166 dominated by *Agrostis capillaris* L., *Festuca rubra* L., *Anthoxanthum odoratum* L., *Ranunculus*  
167 *montanus* Willd., *Leontodon hispidus* L., *Trifolium repens* L. and *T. pratense* L. (Bahn et al., 2009;  
168 Schmitt et al., 2010).

169 The forested observation plot is located close to the tree line at 1960 m a.s.l. on a slope with an  
170 inclination of 20-35°. It is dominated by the two common tree species European larch (*Larix decidua*  
171 Mill.) and Norway spruce (*Picea abies* (L.) H.Karst.). In former years, the plot was a pasture and it was  
172 reforested in the 1980s (Oberleitner et al., 2022).

173 Both observation plots are equipped with micrometeorological stations, soil environment monitoring  
174 (soil moisture, soil temperature), and soil CO<sub>2</sub> devices. At both observation plots, we measured soil

175 CO<sub>2</sub> fluxes with automated chambers during the summer. The forest plot is additionally equipped  
176 with tree dendrometers and tree sapflow sensors. In the grassland, land use and drought related  
177 carbon cycle research was carried out over the last two decades (Fuchslueger et al., 2014; Hasibeder  
178 et al., 2015; Ingrisch et al., 2020; Ingrisch et al., 2018). Research using the forest plot started only  
179 recently (Oberleitner et al., 2022).

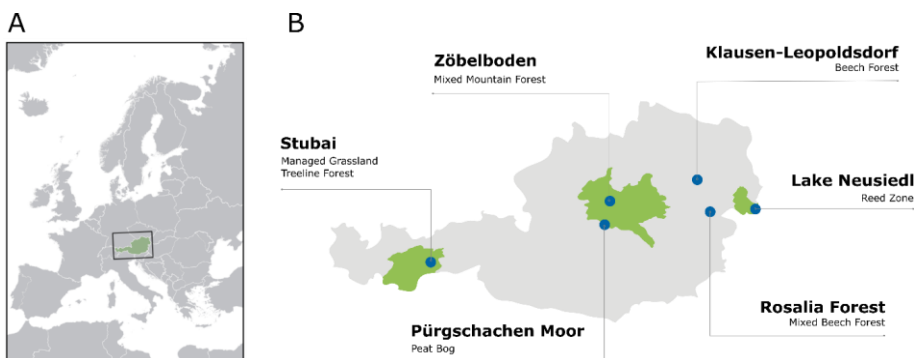
## 180 2.6. Zöbelboden (mixed Beech forest)

181 The site LTER Zöbelboden is located in the National Park Kalkalpen in the Northern Limestone Alps,  
182 Austria (Figure 1). The measurements were established in 1992 as part of the International  
183 Cooperative Programme on Integrated Monitoring of Air Pollution Effects on Ecosystems (ICP IM)  
184 covering a 90 ha catchment with an elevation range of 550 to 956 m a.s.l. (Table 1). The main  
185 underlying rock type is Norian dolomite (*Hauptdolomit*), partly overlain by limestone (*Plattenkalk*).  
186 According to long-term meteorological measurements (1993-2022), mean annual air temperature and  
187 precipitation are 8.2 °C and 1645 mm, respectively. Maximum precipitation occurs in summer and  
188 snowfall usually between December and April.

189 The data presented here was measured at the Intensive Plot II situated on a steep (36° on average)  
190 north-westerly exposed slope at 880 m a.s.l. The soils of the plot are lithic and rendzic leptosols  
191 (Working Group WRB 2015). The plot is dominated by beech (*Fagus sylvatica* L.) with intermixed  
192 sycamore (*Acer pseudoplatanus* L.), European ash (*Fraxinus excelsior* L.) and spruce (*Picea abies* (L.)  
193 H.Karst.). Since the year 1995, this plot is equipped with a number of field measurement devices for  
194 long-term monitoring (throughfall deposition, litter fall traps, lysimeters, soil moisture and  
195 temperature sensors, manual dendrometers) and supplemented by other monitoring activities (tree  
196 inventory, needle and leaf chemistry, soil chemistry, etc.; see e.g. Leitner et al., 2020; Kobler et al.,  
197 2019; Dirnböck et al., 2016; Dirnböck et al., 2020). Drought-impacts on carbon allocation in the  
198 forests of the catchment is currently one of the research foci for which long-term observation data  
199 exists (see e.g. Hartl-Meier et al., 2014) as well as experimental plots with rainout shelters.

200 The instruments and data included here are soil respiration automated chambers, soil water  
201 potential and temperature sensors as well as automated dendrometers. The meteorological data  
202 stems from a station in close proximity at the plateau at 890 m a.s.l.. The site is also equipped with  
203 an Eddy covariance tower, but this data will be published elsewhere.

204



205

206 Figure 1. Location of sites in A) Europe and B) Austria. Blue dots indicate the sites; green areas are long-term socio-  
 207 ecological research platforms (LTSER) (from left to right: LTSER Tyrolian Alps, LTSER Eisenwurzen, and LTSER Lake Neusiedl)  
 208 within the LTER Austria network. [For site information concerning altitude and climate see Table 1.](#)

209 Table 1. [Ecosystem type, altitude, climate and metadata links](#) of the sites and observation plots. Geographic boundaries,  
 210 linked data sets, etc. can be found in the site and dataset registry system DEIMS-SDR.

Site	Ecosystem type	Altitude (m a.s.l.)	Annual temperature (°C)	Annual precipitation (mm)	Site and observation plot	DEIMS.iD
Rosalia Forest Demonstration Centre (ROS)	Mixed beech forest	600	6.5	796	Site	<a href="https://deims.org/77c127c4-2ebe-453b-b5af-61858ff02e31">https://deims.org/77c127c4-2ebe-453b-b5af-61858ff02e31</a>
					Heuberg Meteorological Station	<a href="https://deims.org/locations/44854b32-64c3-4c9d-9aec-9b0b74f8ac70">https://deims.org/locations/44854b32-64c3-4c9d-9aec-9b0b74f8ac70</a>
					Kuhwald Meteorological Station	<a href="https://deims.org/locations/1225d57e-02da-47fd-9760-ab39d64999ef">https://deims.org/locations/1225d57e-02da-47fd-9760-ab39d64999ef</a>
					Mehlbeerleiten Meteorological Station	<a href="https://deims.org/locations/0becf0ce-98d7-4f64-a074-f89046083e5e">https://deims.org/locations/0becf0ce-98d7-4f64-a074-f89046083e5e</a>
Klausen-Leopoldsdorf (KLL)	Beech forest	520	8	801	Site	<a href="https://deims.org/bb472a51-f85f-4de0-8358-f21ecbe2a102">https://deims.org/bb472a51-f85f-4de0-8358-f21ecbe2a102</a>
					Measuring station	<a href="https://deims.org/locations/d5cba3ce-7489-46d1-8d97-61641ffb5758">https://deims.org/locations/d5cba3ce-7489-46d1-8d97-61641ffb5758</a>
Lake Neusiedl (NSS)	reed zone	120	11.5	576	Same as site	<a href="https://deims.org/locations/4234987b-9031-4332-9bdd-f869d503ac51">https://deims.org/locations/4234987b-9031-4332-9bdd-f869d503ac51</a>
Pürgschachen Moor (PUE)	peat bog	632	8.2	1233	Same as site	<a href="https://deims.org/locations/ab2d021b-f318-487a-a85b-ab34566e4c02">https://deims.org/locations/ab2d021b-f318-487a-a85b-ab34566e4c02</a>
Stubai (KAS)	grassland	1830	3	1100	Site	<a href="https://deims.org/324f92a3-5940-4790-9738-5aa21992511c">https://deims.org/324f92a3-5940-4790-9738-5aa21992511c</a>
	Treeline forest	1960	3	1100	Kaserstattalm meadow Kaserstattalm forest	<a href="https://deims.org/locations/cf7843b7-32d6-44e9-ba82-9a8d915036a7">https://deims.org/locations/cf7843b7-32d6-44e9-ba82-9a8d915036a7</a> <a href="https://deims.org/locations/af2afdad-d6fb-4580-b6e3-be7d07b56f8e">https://deims.org/locations/af2afdad-d6fb-4580-b6e3-be7d07b56f8e</a>
Zöbelbo	Mixed forest	880	8.2	1645	Site	<a href="https://deims.org/8eda49e9-1f4e-4f3e-b58e-e0bb25dc32a6">https://deims.org/8eda49e9-1f4e-4f3e-b58e-e0bb25dc32a6</a>

211

## 212 3. Dataset description, measuring methods, QA/QC

213 We followed routine quality assurance (QA) and quality control (QC) procedures to ensure  
 214 functionality of the sensors and data quality comprising remote function control, on-site check of  
 215 sensors and cables, regular sensor calibration, data checks through different quality assurance  
 216 procedures (e.g. exceedance of thresholds, outlier detection, deviations from other measurements),  
 217 and data quality flagging.

## 218 3.1. Meteorology, soil temperature and soil moisture

219 All meteorological stations are located within the boundaries of the respective sites except for  
 220 Klausen-Leopoldsdorf, where the station is at a distance of 2.7 km from the site. Meteorological  
 221 measurements in the wetland sites were implemented next to the Eddy Covariance tower. In  
 222 addition to the routine data checks, we compared the measurements with nearby stations where  
 223 appropriate. Meteorological measurements were detected in a one-minute-interval and averaged  
 224 over half-hour periods while rain data was summed. The measurements include air temperature,  
 225 precipitation, relative humidity, wind speed and direction, air pressure, and several radiation  
 226 variables (at least global radiation, but also short- and longwave radiation, photosynthetic active  
 227 radiation, etc.).

228 *Table 2. Meteorological parameters measured at the six sites during 2019-2021. Availability is indicated using grey boxes.*  
 229 *For parameter names we refer to the thesaurus at <http://vocabs.lter-europe.net/EnvThes/>*

230

## Meteorological parameters

	KAS	KLL	NSS	PUE	ROS	ZOE
air relative humidity						
air pressure						
air temperature						
precipitation amount						
global radiation irradiance						
net radiation irradiance						
photosynthetically active radiation						
sunshine duration						
wind direction						
wind speed						

231

232 We used different types of soil temperature and soil moisture or soil water potential sensors,  
 233 respectively (PT100 or thermoelements for soil temperature, TDR or FDR-sensors for soil moisture, and  
 234 soil water potential sensors). Before we buried the soil temperature or soil moisture and soil water  
 235 potential sensors into the soil, they had been calibrated or at least tested for consistency. Mostly, we  
 236 used gravimetric samples to calibrate the TDR and FDR soil moisture sensors. At Zöbelboden, where  
 237 stony, organic rich soils occur, we corrected the TDR values using water potential sensor data installed  
 238 in the same soil profiles together with soil water retention functions derived from undisturbed soil  
 239 cores. In addition to the regular QC procedures, we checked the data for consistency of the values

240 across sensors (e.g. along the soil profiles) and compared them with other measurements (air  
 241 temperature and precipitation). Half-hourly to hourly values are presented.

### 242 3.2. Carbon fluxes

#### 243 3.2.1. Soil CO<sub>2</sub> efflux

244 We measured soil CO<sub>2</sub> efflux at five of the seven observation plots. The automated soil CO<sub>2</sub> respiration  
 245 measurement systems are capable of operating autonomously during the snow-free periods. The  
 246 measurement chambers and measurement systems collected air from the chamber headspace  
 247 continuously to determine the exchange of CO<sub>2</sub> between soil and atmosphere at the observation plots.  
 248 In all sites, we used non-steady state, non through-flow chambers (Pumpanen et al., 2004). The  
 249 chambers at each site measured consecutively every half-hour to hour. In addition to the automated  
 250 systems, manual flux measurements were also performed which served to validate the automated  
 251 measurement systems. Table 3 provides detailed information on the measurement systems used at  
 252 the sites.

253 Two different automated chamber systems were used: a LI-COR System and custom-made chambers  
 254 in combination with LI-COR trace gas analysers (Table 3). The custom-made soil chambers are  
 255 equipped with a fan and a thermometer. The controlling unit and the gas analyzer (either a CH<sub>4</sub>/CO<sub>2</sub>  
 256 LI-COR 7810, a LI-COR 840, or a LI-COR 8100A, LI-COR Biosciences, USA) are located in already  
 257 existing measurement containers. Remote access to the devices allows for checking plausibility of the  
 258 data and chamber leakage in real time. We visited the instruments at weekly to monthly intervals,  
 259 with maintenance and supervision works including a check of the tightness of the gas lines,  
 260 connections and chamber lids, the correct closing and opening of the chambers and the functioning  
 261 of ventilation fans inside the chambers, ingrowth of plants, and the gas analyser. The gas analysers  
 262 were calibrated once a year in the laboratory with calibration gases. We de-installed and serviced the  
 263 chambers during winter but frames stayed permanently on site to avoid disturbance of the soil.

264 At Klausen-Leopoldsdorf, the gas fluxes of readings were determined using the R package "gasfluxes"  
 265 (Fuss, 2020). At Rosalia, a custom-made Python script was used. Zöbelboden and Kaserstattalm  
 266 process the data with SoilFlux Pro Software (LI-COR Biosciences, 2019). We used the R<sup>2</sup> of the fitted  
 267 empirical models to select valid data. We refer to Table 3 and the metadata published with the data  
 268 for the detailed specifications.

269 *Table 3. Specifications of the different soil CO<sub>2</sub> flux systems following the standard of (Bond-Lamberty et al., 2021).*

Field Name	Description	Unit	Klausen- Leopoldsdorf	Stubai grassland	Rosalia	Zöbelboden
System			auto	Auto	auto	auto
GHG chambers			Custom-made (n=12)	LI-8100-104 (n=4)	Custom- made (n=12)	LI-8100-104 (n=6)
INSTRUMENT	Measurement instrument model		LI-COR LI-7810	LI-8100A	LI-840	LI-8100A
MSMT_VAR	Type of flux measured		Soil respiration (Rs)			
AREA	Soil surface measurement area	cm <sup>2</sup>	2500	317.8	2500	317.8



VOLUME	Volume of measurement chamber	cm <sup>3</sup>	37500	4076.1	37500	4076.1
V/A	Volume/Area ratio	cm	15	12.83	15	12.83
COLLAR_DEPTH	Depth of collar insertion	cm	5	2	10	2
OPAQUE	Opaque chamber		no	Yes	no	yes
chamber system	static chamber - closed or open		non-steady state, non through-flow chambers			
closing time	closing time of chamber (=time used for flux calculation)	sec	175	depending on year	1620	210
PLANTS_REMOVED	Plants removed from inside the collar		no, but hardly any	Yes	no, but hardly any	no plants
flow_rate	sample flow rate through tubing	l min <sup>-1</sup>	1	1 to 2	0.25	1.7
FAN	Mixing fan in chamber?		yes	No	yes	no
CRVFIT_CO2	Flux computation method ("Lin" or "Exp" for linear and exponential, others)		linear	automated <sup>1</sup>	Lin/HMR <sup>2</sup>	Automated <sup>1</sup>
R2_CO2	R <sup>2</sup> of flux computation	fraction	0.90	0.95	0.95	0.99
Calculation of flux			R Package gasfluxes	LI-COR Soilflux Pro	custom-made python script	LI-COR Soilflux Pro

<sup>1</sup> "Exp" in the data indicates that the exponential fit was better than the linear fit (Exp\_SSN < Lin\_SSN). "Lin" indicates that the linear fit was better after the maximum number of iterations; the non-linear coefficients have therefore been derived from the linear fit.

<sup>2</sup> Hutchinson and Mosier (1981)

### 270 3.2.2. Eddy Covariance measurements at wetland sites

271 In both wetland sites, the Pürgschachen Moor and Lake Neusiedl, fully equipped Eddy-Covariance  
272 systems are in place. Wind speed and direction were measured using a three-axis ultrasonic  
273 anemometer (WindMaster Pro, Gill Instruments, Lymington, UK). CO<sub>2</sub> and H<sub>2</sub>O mixing ratios were  
274 measured using the closed-path infrared gas analyser LI-7200 while CH<sub>4</sub> was detected with the open  
275 path gas analyser LI-7700 (both LI-COR Inc, Lincoln, USA). The measurements were performed with a  
276 sampling rate of 10 Hz. We installed the devices at a vegetation dependent height, 3.05 m above  
277 ground in the Pürgschachen Moor and in the reed belt of Lake Neusiedl 8.6 m, respectively. The Eddy  
278 Covariance devices were checked daily via remote access, calibrated once a year in the lab, and  
279 monthly in the field.

280 The EC data contains half-hour eddy covariance flux measurements for CO<sub>2</sub>, CH<sub>4</sub> and water vapor. We  
 281 calculated the fluxes with the EddyPro® Software package in the Express mode with default settings  
 282 (double rotation, block averaging, covariance maximization, etc.) as part of the SmartFlux® 2 System,  
 283 providing fully corrected and valid fluxes with quality flags ranging from 0-2. The final flags are based  
 284 on a combination of partial flags accounting for steady state and turbulent conditions. Only fluxes  
 285 flagged with 0 (best quality fluxes) or 1 (fluxes suitable for general analysis such as annual budgets)  
 286 are shown in the data. Gaps in the data-set result from missing micro-meteorological conditions,  
 287 from data cleaning due to the quality flags or from power breakdowns.

### 288 3.3. Radial tree stem growth at forest sites

289 Zöbelboden, Klausen-Leopoldsdorf, Rosalia used the DR26 sensor (EMS, Brno, Czech Republic), Stubai  
 290 used Ecomatic DC2 (Germany) for registering the radial stem increment in a 15 minutes to 30 minutes  
 291 interval. Maintenance involved avoiding any shift of the sensor during the operation. Concerning data  
 292 quality and control methods the Mini32 software (EMS, Brno, Czech Republic), includes graphical  
 293 features to process the measured stem increment data. Data processing comprises outlier detection  
 294 by visual assessment based on expert knowledge. Ecomatic raw data was treated with custom-made  
 295 R scripts. In both cases, unrealistic values beyond the slowly increasing linear growth rates were  
 296 visually assessed and deleted.

### 297 4. Data file structure

298 We used the eLTER Data specification, which is available on Zenodo  
 299 ([www.doi.org/10.5281/zenodo.6373409](http://www.doi.org/10.5281/zenodo.6373409)). Apart from the data files, the measurement locations  
 300 (Station files) and the sensors (methods) are included.

301

### 302 5. Data validation

303 *Table 4. Comparison between long-term (1980-2010) meteorological drought (SPEI - Standardized Precipitation*  
 304 *Evapotranspiration Index) and the measurement years during the growing season (May-September). Significant differences*  
 305 *between these years and the long-term averages are shown: \*\*\* p<0.001; \*\* p<0.005; \* p<0.01 according to a Mann-*  
 306 *Whitney U Test. SPEI was calculated using a 30 days window in a daily resolution using gridded data:*  
 307 *<https://data.hub.geosphere.at/dataset/winfore-v2-1d-1km> (Haslinger & Bartsch, 2016). Negative values indicate dry years.*

Site Code	1980-2010	2019	2020	2021
ZOE	-0.05±0.94	-0.49±1.22***	0.42±0.69***	-0.02±1.39
KLL	-0.01±0.99	-0.23±1.22	0.39±0.86***	-0.02±1.24
KAS	-0.05±0.97	-0.1±1.17	-0.12±0.78	0.23±1.07**
PUE	-0.02±0.94	-0.54±1.26***	0.17±0.59**	-0.12±1.2
NSS	-0.01±1	-0.22±1.13*	0.2±0.87**	-0.2±1.34
ROS	-0.03±0.98	-0.34±1.01***	0.17±0.83	-0.28±1.08

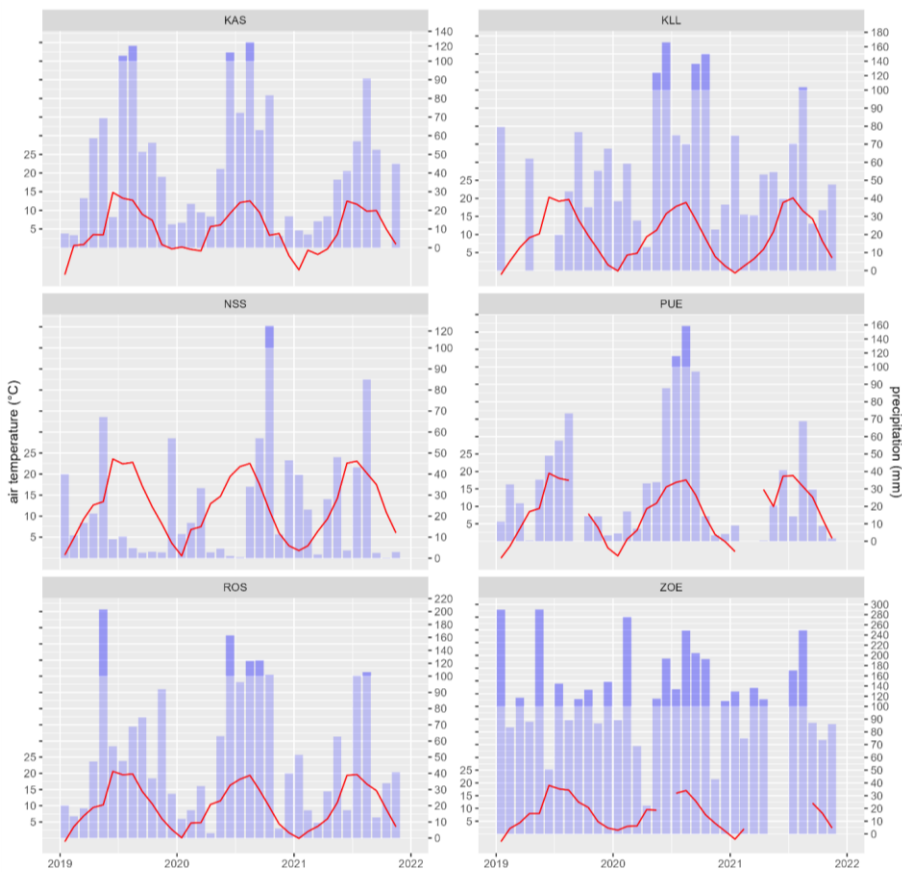
308

309 We used gridded SPEI (Standardized Precipitation Evapotranspiration Index) from the Austrian  
 310 Meteorological Service (<https://data.hub.geosphere.at/dataset/winfore-v2-1d-1km>; Haslinger &  
 311 Bartsch (2016)) to compare the long-term average water availability during the growing season  
 312 (1980-2010; May to September) with those occurring in the measurement years (Table 4). The  
 313 advantage of the SPEI is that it accounts for precipitation and temperature via evapotranspiration  
 314 and integrates over a given temporal window (we used 30 days) (Vicente-Serrano et al. 2010). Note,  
 315 that gridded SPEI data set is based on meteorological data for the period 1960 to 2021. Accordingly,  
 316 the 2021 was closest to the long-term average, the year 2020 was a particularly wet year, and the

317 year 2019 was drier than the average. However, there were differences between the sites:  
 318 particularly the mountain station in the Tyrolian Alps (KAS) did not experience significant deviations  
 319 in SPEI as compared to the long-term average apart from a wet growing season in 2021. The SPEI at  
 320 the site in the Viennese Forest (KLL) does not indicate that in 2019, the growth period was  
 321 particularly dry.

322 The monthly precipitation and temperature patterns are shown in [Figure 2](#), and soil water  
 323 content and soil temperatures in [Figure 3](#) and Figure 5. Differences in the seasonal  
 324 precipitation patterns between the measurement years vary a lot between sites. In sum, lower  
 325 precipitation occurred in 2019 and 2021 than in 2020 in all sites. The mean annual temperature  
 326 maxima (90 percentile) were between 0.3 °C (KAS) and 2.3 °C (ZOE) higher in the year 2019 than in  
 327 2020. These differences were lower when comparing the year 2021 with 2019 ( $\leq 0.6$  °C). In  
 328 accordance with SPEI, precipitation and temperature, soil water content showed the lowest values  
 329 during the years 2019 followed by the year 2021, and soil temperature were higher during these  
 330 years ([Figure 3](#)).

331



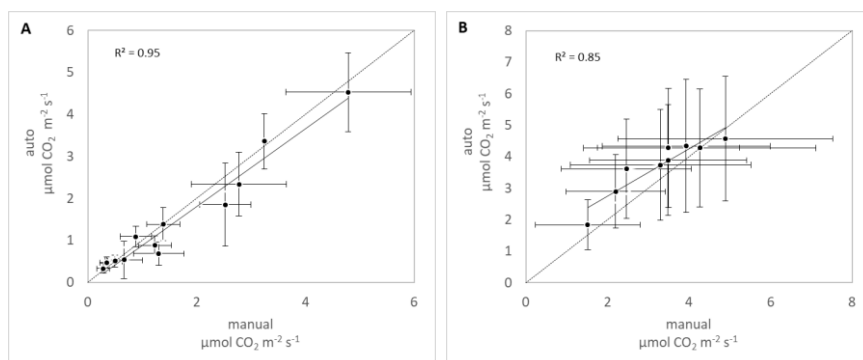
332

333 *Figure 2. Monthly mean air temperature (red line) and monthly precipitation sums (blue bars; different scale > 100 mm) at*  
334 *the six regions.*

335 We measured soil CO<sub>2</sub> respiration at four sites (Figure 4). The complexity of automated chamber  
336 measurements resulted in some data gaps: at KAS and ROS during the years 2019 and 2021  
337 respectively; at KLL and ZOE, the respiration data covers most of the snow-free period (see Table 2).

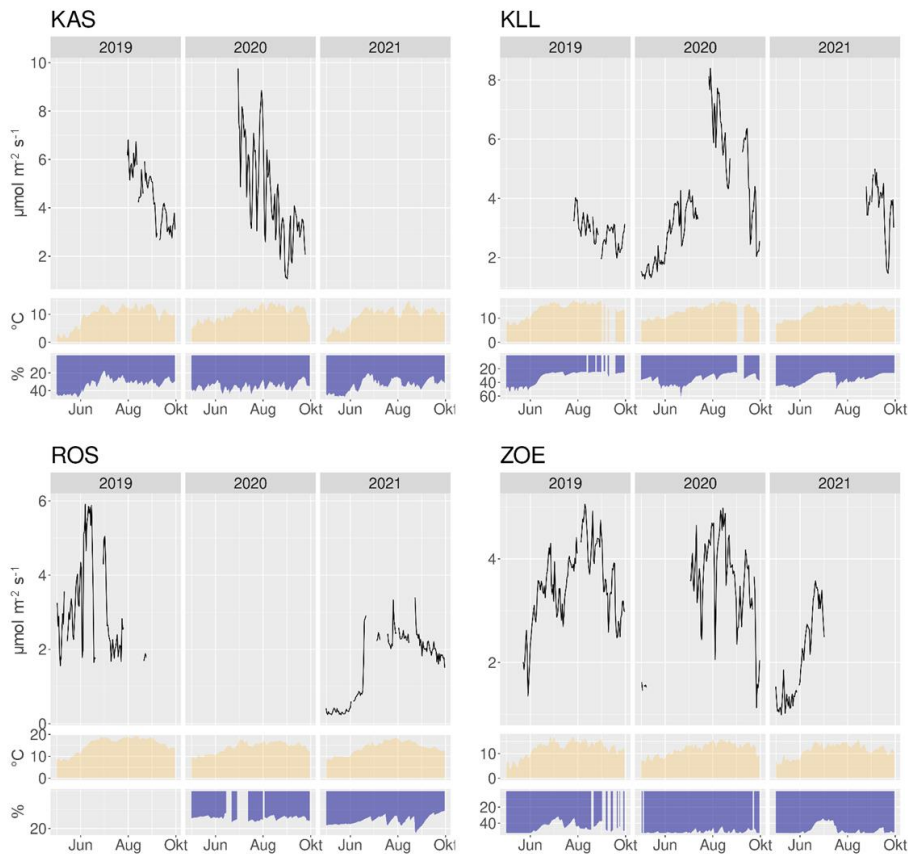
338 At Klausen-Leopoldsdorf (KLL) and Zöbelboden (ZOE), we compared the automatically measured soil  
339 CO<sub>2</sub> flux rates with manual measurements. For both sites, we used a portable infrared gas analyzer  
340 (EGM-4) connected to a manual soil respiration chamber (SRC) (PP Systems International Inc.,  
341 Amesbury, MA, USA). The two measurement sites were equipped with permanently installed collars  
342 (KLL: randomly distributed within the site in immediate vicinity of the automated chambers (n = 12);  
343 area = 284 cm<sup>2</sup> and 2 cm insertion depth; Zöbelboden: regular grid covering the entire plot (n = 30),  
344 area = 78 cm<sup>2</sup> and 1.5 cm insertion depth). The chamber closure time was 60 and 100 seconds in KLL  
345 and ZOE, respectively. Manual measurements took place in monthly intervals from Oct. 2019-Jun.  
346 2020 at Klausen-Leopoldsdorf and from Jun. 2019 until Oct. 2019 (monthly interval) and in July 2020  
347 (diurnal variation) at ZOE. Soil respiration (Rs) was calculated automatically by fitting a linear (KLL) or  
348 quadratic function (ZOE; quadratic fit for flow rates > 0.2 ppm s<sup>-1</sup>, otherwise a linear fit was used) to  
349 the increasing CO<sub>2</sub> headspace concentration.

350 The mean CO<sub>2</sub> fluxes of the automated chambers correlated well with the manually measured fluxes  
351 during the measurement campaigns (Figure 3). At KLL, the R<sup>2</sup> was 0.95 (p-value < 0.05, t-  
352 test), at ZOE it was 0.85 (p-value < 0.05, t-test). In both sites, neither the intercept nor the slope was  
353 significantly different from 0 (p-value > 0.2, t-test) and 1 (p-value > 0.49, t-test), respectively. At ZOE,  
354 the spatial flux variation was much higher than at KLL (Figure 3A and 3B). This reflects the  
355 heterogeneity of the soil conditions (shallow rendzic leptosols with interspersed fine-scale patches of  
356 deeper soils), the canopy gaps (with lower root density), and the uneven distribution of litter due to  
357 the steep slope at the plot (Kobler et al. 2019). This heterogeneity is more effectively captured in  
358 the manual measurement (n=30) than by the automated chambers (n=6). In summary, we conclude  
359 that the spatial variation in CO<sub>2</sub> fluxes was higher at both sites than the difference in fluxes caused by  
360 the measurement devices (Figure 3).



361  
362 *Figure 3. Comparison of automated and manual soil CO<sub>2</sub> fluxes at A) Klausen-Leopoldsdorf and B) Zöbelboden. See Table 3*  
363 *for the specification of automated chamber data. Error bars indicate spatial variation (standard deviations).*

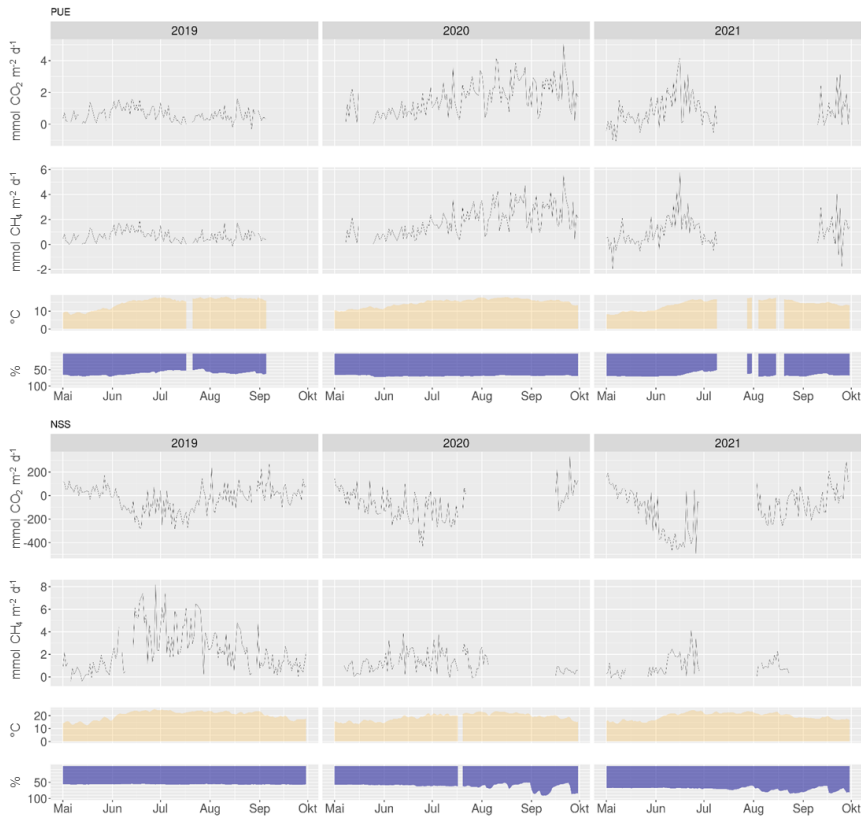
364



365

366 *Figure 4. Soil CO<sub>2</sub> respiration (mean of all chambers), soil temperature (mean of sensors in 5-15 cm depth) and soil water*  
 367 *content (mean of sensors in 5-15 cm depth) in the forested sites Kaserstattalm forest (KAS), Klausen-Leopoldsdorf (KLL),*  
 368 *Rosalia (ROS), and Zöbelboden (ZOE).*

369 Soil CO<sub>2</sub> fluxes are temperature dependent, thus follow the seasonal changes in soil temperature  
 370 (Figure 4). Their additional limitation through soil water availability for plant metabolism and  
 371 microbial activity is usually much less pronounced in these temperate zone ecosystems (Bahn et al.  
 372 2008; Chen et al. 2014). For detailed interpretation of the CO<sub>2</sub> respiration fluxes and their limiting  
 373 factors, we refer to the citations listed in the site description chapter. Drollinger et al. (2019)  
 374 provides interpretations of the patterns of CO<sub>2</sub> and CH<sub>4</sub> fluxes, measured using Eddy covariance  
 375 techniques at the bog site Pürgschachen Moor (PUE), and likewise, Baur et al. (2024), for the reed  
 376 belt of Neusiedler See (NSS). Stem growth limitations can, on the other hand, be closely related to  
 377 soil water content, particularly at sites with relatively low precipitation such as Klausen-Leopoldsdorf  
 378 (KLL) (Figure 6). For an in-depth study of drought related effects on tree growth at the treeline forest  
 379 at Kasterstattalm (KAS), we refer to Oberleitner et al. (2022).



380

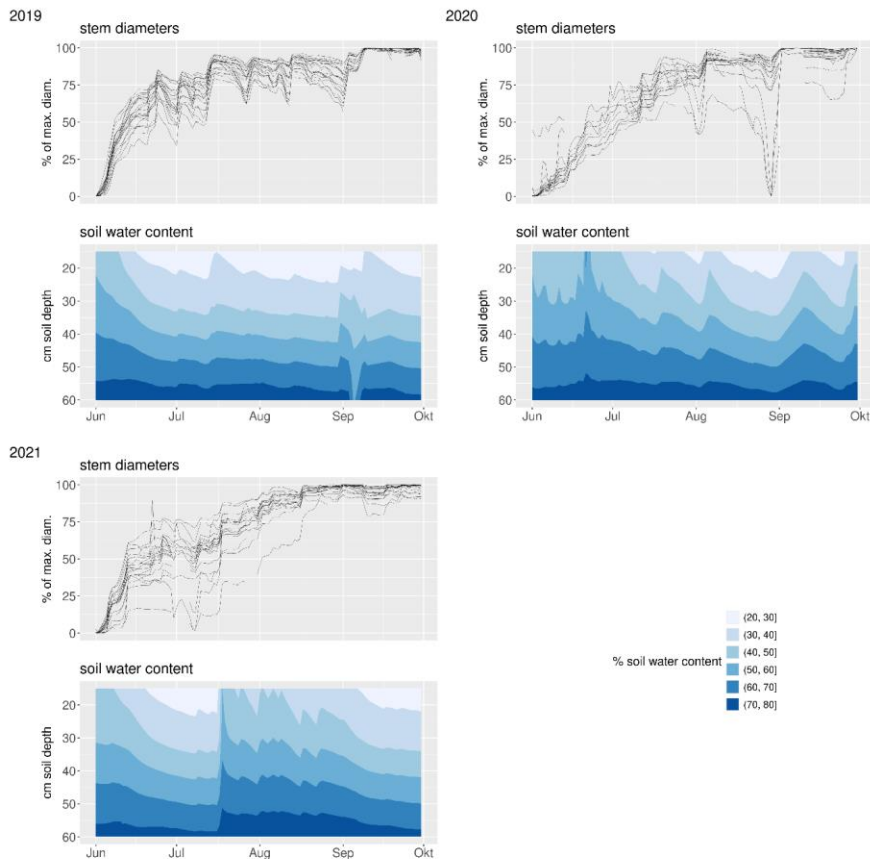
381

382

383

384

Figure 5. CO<sub>2</sub> and CH<sub>4</sub> fluxes in the sites Pürgschachen Moor peat bog (PUE), Lake Neusiedl reed zone (NLL) and Pürgschachen Moor peat bog (PUE)-Lake Neusiedl reed zone (NLL) as well as temperature and soil water content and temperature



385  
 386 *Figure 6.* Relative stem diameters and soil moisture at the site Klausen-Leopoldsdorf (KLL) during the ~~dry~~ years 2019 (A),  
 387 2020 (B), and 2021 (C). Stem diameter values were scaled to an annual amplitude of 100.

388

389 6. Discussion

390 We provide baseline ecosystem data related to the carbon cycle and capture naturally occurring ECEs  
 391 across various ecosystem types typical for Austria and other regions of Central Europe. Such data sets  
 392 are scarce because the measurements are demanding in terms of maintenance and funding.  
 393 Automated soil respiration data in high temporal resolution, as we report it here, is rare too owing to  
 394 a lack of dedicated monitoring or research infrastructures (Bond-Lamberty et al., 2021). However,  
 395 soil CO<sub>2</sub> respiration constitutes the second-largest flux in the global carbon cycle, hence is key in  
 396 estimating ecosystem response to ECEs (Bond-Lamberty and Thomson, 2010). In addition, we  
 397 provide soil temperature and moisture measurements in the same resolution, being key variables  
 398 determining soil respiration (Pumpanen et al., 2015). High-resolution measurements of tree stem  
 399 circumference have been developed as complementary data to relate drought stress with changes in  
 400 carbon allocation in trees (Zweifel, 2016; Zweifel et al., 2021). The microclimatic, soil, and tree

401 physiological data is complemented by CO<sub>2</sub> and CH<sub>4</sub> fluxes between the vegetation and the  
402 atmosphere measured with Eddy covariance techniques of the two wetland sites.

403 Our data is particularly useful for drought-related research. Triggered by the pan-European drought  
404 of 2003 (Ciais et al., 2005), a key scientific question has been how droughts affect greenhouse gas  
405 sinks and sources in ecosystems (Rödenbeck et al., 2020; Reichstein et al., 2013; Anderegg et al.,  
406 2020). Droughts usually reduce soil respiration due to the decrease in autotrophic respiration but  
407 also because soil microbial activity drops due to water limitation (Grünzweig et al., 2022).  
408 Furthermore, rewetting can result in pulses of high soil respiration (Borken and Matzner, 2009).  
409 Drought effects on the ecosystem C cycle can persist for years (Kannenbergh et al., 2020; Müller and  
410 Bahn, 2022) and novel approaches are being developed for assimilating high-resolution data for  
411 understanding and quantifying such legacies (Yu et al., 2022; Fu et al., 2020). In this context, the  
412 availability of long-term, high-resolution measurements of key ecosystem parameters is key for  
413 understanding and quantifying the effects of recurrent droughts (Oberleitner et al., 2022). While the  
414 three-year data with the usual measurement gaps occurring in field campaigns in rather difficult  
415 terrain can only to some extent capture aspects of drought related effects, it represents a valuable  
416 baseline.

417

418 The sites presented here are currently being upgraded towards their implementation in the  
419 European Research Infrastructure for Integrated European Long-Term Ecosystem, critical zone and  
420 socio-ecological Research (eLTER RI), together with another ~200 sites in Europe (Mirtl et al., 2018).  
421 Climate change impacts on ecosystem processes including the carbon cycle are among the targeted  
422 research areas the eLTER RI will focus on. The measurements resulting in the data presented here  
423 will continue in future under the umbrella of eLTER RI. Compiling longer-term data series depends  
424 upon the availability of already validated data sets - as it is presented here - before the RI is being  
425 operational. Furthermore, long-term ecosystem observations already exist in these sites with regard  
426 to water and nitrogen cycle allowing for a contextual interpretation of the trends seen in C related  
427 parameters.

428 Combining several research and monitoring activities at already heavily instrumented sites not only  
429 saves money but widens the data analyses portfolio (Futter et al. 2023; Kulmala 2018). Even though  
430 we provide Eddy covariance data for two of our sites, Austria is not part of the Integrated Carbon  
431 Observation System (ICOS). A combination of data capturing long-term boundary layer exchange of C  
432 together with soil C fluxes, microclimate, and, in forests, tree physiological data obviously holds great  
433 potential (Zweifel et al., 2023; Ramonet et al., 2020). Hence, using the sites simultaneously for other  
434 research infrastructures, such as ICOS, providing high-quality Eddy covariance measurements would  
435 obviously be ideal. The more so because European Research Infrastructures follow the FAIR data  
436 principles to make data Findable, Accessible, Interoperable and Reusable (Wilkinson et al., 2016).

437 While the eLTER RI data infrastructure is still under development, we comply with the standards  
438 already implemented. We used DEIMS-SDR (<https://deims.org/>) as the catalogue documenting the  
439 sites (Wohner et al., 2019; Wohner et al., 2022). It issues persistent identifiers for sites (see Table 1)  
440 that allow to uniquely identify sites across research projects and networks. Tools are being  
441 developed to query available information about sites programmatically (Oggioni et al., 2023;  
442 Wohner, 2023) providing contextual ecosystem information.

443

444 7. Data availability



445 7.1 Data access

446 The data and metadata is accessible at B2SHARE (<https://b2share.eudat.eu/>), a service provided by  
 447 the EUDAT Collaborative Data Infrastructure. DOIs of the datasets are listed in Table 5. The site  
 448 metadata in DEIMS-SDR (Table 1) is part of the data metadata so that site information can easily be  
 449 accessed. In chapter8, we provide a jupyter notebook to download and merge the single datasets,  
 450 and to visualize parameters.

451 *Table 5. Dataset DOIs*

Site	Dataset	DOI	Reference
Klausen- Leopoldsdorf	Meteorology	<a href="https://doi.org/10.23728/b2share.8f872a37513c4768b16ce755eca4bb57">https://doi.org/10.23728/b2share.8f872a37513c4768b16ce755eca4bb57</a>	(Gartner et al., 2024a)
	Soil climate	<a href="https://doi.org/10.23728/b2share.8d49c0b557f1455a9e66689e035b8cce">https://doi.org/10.23728/b2share.8d49c0b557f1455a9e66689e035b8cce</a>	(Gartner et al., 2024b)
	Soil CO <sub>2</sub> respiration	<a href="https://doi.org/10.23728/b2share.5286bd1bc6aa491f874b9bb12d1c5673">https://doi.org/10.23728/b2share.5286bd1bc6aa491f874b9bb12d1c5673</a>	(Kitzler and Hofbauer, 2024)
	Stem increment	<a href="https://doi.org/10.23728/b2share.68d84a913f0c4875be5c680ad4d6959e">https://doi.org/10.23728/b2share.68d84a913f0c4875be5c680ad4d6959e</a>	(Gartner and Gollobich, 2024)
Rosalia Forest Demonstration Centre	Meteorology	<a href="https://doi.org/10.23728/b2share.96c52c247eb846deb2a3ec5e2c27b4f1">https://doi.org/10.23728/b2share.96c52c247eb846deb2a3ec5e2c27b4f1</a>	(Diaz-Pines, 2024a)
	Soil climate	<a href="https://doi.org/10.23728/b2share.c68143fc11224c44ae5529bd6a35a76d">https://doi.org/10.23728/b2share.c68143fc11224c44ae5529bd6a35a76d</a>	(Diaz-Pines, 2024c)
	Soil CO <sub>2</sub> respiration	<a href="https://doi.org/10.23728/b2share.d167e727abe947abbc8efc04057557f6">https://doi.org/10.23728/b2share.d167e727abe947abbc8efc04057557f6</a>	(Diaz-Pines, 2024b)
	Stem increment	<a href="https://doi.org/10.23728/b2share.d0d185f1eb184ae48f6d06ea9aa8dbdf">https://doi.org/10.23728/b2share.d0d185f1eb184ae48f6d06ea9aa8dbdf</a>	(Diaz-Pines, 2024d)
Zöbelboden	Meteorology	<a href="https://doi.org/10.23728/b2share.762e665273234b129d09ef017416bcfb">https://doi.org/10.23728/b2share.762e665273234b129d09ef017416bcfb</a>	(Kobler et al., 2024a)
	Soil climate	<a href="https://doi.org/10.23728/b2share.46e19191ce9c427d90f48ce38f56a0e1">https://doi.org/10.23728/b2share.46e19191ce9c427d90f48ce38f56a0e1</a>	(Kobler et al., 2024c)
	Soil CO <sub>2</sub> respiration	<a href="https://doi.org/10.23728/b2share.4f44006b932142e68981106a016f1f56">https://doi.org/10.23728/b2share.4f44006b932142e68981106a016f1f56</a>	(Kobler et al., 2024b)
	Stem increment	<a href="https://doi.org/10.23728/b2share.2de5b37a0cad4f82a19f477531d6af24">https://doi.org/10.23728/b2share.2de5b37a0cad4f82a19f477531d6af24</a>	(Pröll et al., 2024)
Stubai - Kaserstattalm	Meteorology	<a href="https://doi.org/10.23728/b2share.77462914dc0b43cb8c24a967e6851665">https://doi.org/10.23728/b2share.77462914dc0b43cb8c24a967e6851665</a>	(Ingrisch and Bahn, 2024c)
	Soil climate	<a href="https://doi.org/10.23728/b2share.026d76094e8f4512b09b35b7a0d2a9d7">https://doi.org/10.23728/b2share.026d76094e8f4512b09b35b7a0d2a9d7</a>	(Ingrisch and Bahn, 2024d)
	Soil CO <sub>2</sub> respiration	<a href="https://doi.org/10.23728/b2share.cfe8c7ad1965433484650ea9026512ca">https://doi.org/10.23728/b2share.cfe8c7ad1965433484650ea9026512ca</a>	(Ingrisch and Bahn, 2024a)
	Stem increment	<a href="https://doi.org/10.23728/b2share.0e3eed54ff30418f8720806b5f05cca9">https://doi.org/10.23728/b2share.0e3eed54ff30418f8720806b5f05cca9</a>	(Ingrisch and Bahn, 2024b)
Pürgschachen Moor	Meteorology	<a href="https://doi.org/10.23728/b2share.5442510ad03e4968afb4e2108e85a64d">https://doi.org/10.23728/b2share.5442510ad03e4968afb4e2108e85a64d</a>	(Maier and Glatzel, 2024e)
	Soil climate	<a href="https://doi.org/10.23728/b2share.9380364098d14978b876a87517652d62">https://doi.org/10.23728/b2share.9380364098d14978b876a87517652d62</a>	(Maier and Glatzel, 2024f)
	Eddy Covariance	<a href="https://doi.org/10.23728/b2share.4f783e3ff2884abca5c59960db0b7955">https://doi.org/10.23728/b2share.4f783e3ff2884abca5c59960db0b7955</a>	(Maier and Glatzel, 2024d)
Lake Neusiedl	Meteorology	<a href="https://doi.org/10.23728/b2share.f7176c9ee982464f947d2fe9fb8f389d">https://doi.org/10.23728/b2share.f7176c9ee982464f947d2fe9fb8f389d</a>	(Maier and Glatzel, 2024b)
	Soil climate	<a href="https://doi.org/10.23728/b2share.4e6474cd55f9487d97e3d31e83baa530">https://doi.org/10.23728/b2share.4e6474cd55f9487d97e3d31e83baa530</a>	(Maier and Glatzel, 2024c)

Eddy <https://doi.org/10.23728/b2share.b83caca3efe44868a1ed49129b4a576a> (Maier and Glatzel, 2024a)  
Covariance

---

452

## 453 7.2 Data visualization, workflow integration

454 The software stack used to store, import and quality control the provided data is built on PostgreSQL  
455 database with a Post-GIS extension. The database structure is derived from the Time Series  
456 Management (TSM) system developed by the Research Center Jülich (Wohner, C., Dirnböck, T.,  
457 Peterseil, J., Pröll, G., Geiger, S., 2021) and originally deployed during the LTER CWN project but was  
458 repurposed to better fit the needs of the data management and working group. Now, for the import  
459 and quality control of data, a number of Python scripts deployed in a Jupyter environment are used.  
460 This is also includes scripts to visualise the data on the fly in Jupyter.

461

## 462 8. Code availability

463 A Jupyter notebook to access, merge, and visualize the data from all sites is available at  
464 <https://gist.github.com/10/9bbe44a03f12801c6c742202b005db57>.

465

## 466 9. Author contribution

467 DT, BM, DPM, DI, EM, GK, GG, MA, IJ, KB, KJ, MA, PG, VS, ZBS, ZA, and GS designed the  
468 measurements and carried them out. WC, PJ designed and constructed the database. KK, VS, and PG  
469 customized and filled the database. OI developed the Jupyter notebook. DT prepared the manuscript  
470 with contributions from all co-authors.

471

## 472 10. Competing interests

473 The authors declare that they have no conflict of interest.

474

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