Review of "A Copernicus-based evapotranspiration dataset at 100-m spatial resolution over the Mediterranean region" by Bartkowiak et al.

This paper presents an original evapotranspiration (ET) data set at 100 m resolution over 4 basins of the Mediterranean region: Ebro basin in Spain, Po basin in Italy, Herault basin in France and Medjerda basin in Tunisia. The main originality of the data set is its high spatial resolution compared to that of existing ET products classically available at 1 km or coarser resolution. The new 100 m resolution ET data set is derived by automatizing existing codes based on TSEB (Two-Source Energy Balance) model and Sentinel-2 (S2) and Sentinel-3 (S3) remote sensing data. The satellite-derived ET estimates are evaluated with eddy covariance measurements collected at 8 sites with 7 located in Italy and 1 in France and with several land covers (grassland, evergreen broadleaf forest, evergreen needleleaf forest and vineyard). Although the presented data set may be of interest for many different applications over the basins studied, I think that the evaluation strategy must be improved to really demonstrate the better accuracy of the new data set.

We would like to thank the reviewer for revising the manuscript and providing helpful suggestions and comments to improve this work. Below we provide answers to the reviewer’s comments.

I recommend major revisions taking into account the concerns listed below:

1) Title, abstract and conclusion: The extent of the data set is confusing and somehow over-sold as the actual data set does not cover the entire Mediterranean region but only four selected basins within the Mediterranean region. It is true that the algorithms developed by the authors should work over other parts of the Mediterranean region but the paper focuses on the dataset. The authors should be more specific in the title, abstract and conclusion.

Response 1: Thank you very much for this useful comment. Indeed, it is true that our data paper focuses on 100-m evapotranspiration estimation over selected parts of the Mediterranean region. In order to avoid confusion regarding dataset extent, we have modified the title and the Conclusions section. In the abstract we directly indicate our study areas (i.e., Ebro, Po, Herault, and Medjerda basins) where ET is produced.

2) Evaluation of the 100 m resolution evapotranspiration dataset: The evaluation of satellite-derived evapotranspiration estimates is generally sound. However in my opinion it suffers from two major weaknesses. As outlined in the abstract and introduction and other parts of the paper, the rationale for developing a new ET product at high spatial resolution is that common products available at coarser spatial resolution are not sufficient to characterize the very high heterogeneity of land surfaces. The validation strategy of their product should support this key point. This is all the more needed as the ET product relies on the downscaling of 1 km resolution S3 land surface temperature (LST) data from 1 km to 100 m resolution. The evaluation of 100 m satellite-derived ET must be consolidated by estimating the gain in accuracy provided by the use of 100 m resolution remote sensing data, instead of 1 km resolution remote sensing data. One way of achieving this would be to implement PT-TSEB at 1 km resolution at the validation sites and calculate performance metrics as is done for the 100 m resolution dataset. Another drawback of the validation exercise is that it is based on only 8 stations, with 7 located in the same (Po) basin. Readers need to be convinced that this data set is significantly better than other more classical data sets.

The rationale for developing a new ET product at high spatial resolution:
Line 9-10: "existing global products with spatial resolution >=0.5 km are insufficient to capture spatial detail at a local level"
Response 2: We fully agree with your point that the TSEB performance driven by both original Sentinel-3 LST data and its 100-m downscaled product shall be carried out in order to evaluate the effectiveness of thermal sharpening on ET retrievals. In this work, we have exploited data mining sharpener (DMS) of Gao et al. (2012) successfully used in many research studies for estimating high spatial resolution TIR-ET (Anderson et al., 2021; Guzinski and Nieto, 2019; Yang et al., 2021; Guzinski et al., 2023). As presented by Guzinski and Nieto (2019), TSEB-PT driven by downscaled DMS-based surface temperatures is more performant compared to ET estimates driven by original 1-km LST data with around a 13% increase in Pearson correlation coefficient (r) between in-situ ETs and their corresponding modelled observations. Furthermore, the authors of the ESA Sen-ET report estimated evapotranspiration using METRIC, ESVEP, and TSEB-PT algorithms at 11 flux tower sites across different vegetation types and climate zones, and derived the best accuracy scores from the latter model when data mining LST sharpener (either based on Artificial Neural Networks or Decision Trees regressors) was applied (https://www.esa-sen4et.org/downloads/prototype_evaluation_v1.3.pdf). According to the report, the Priestley-Taylor Two-Source Energy Balance of ET was ranked as the most robust approach with consistently lower Root Mean Square Error (RMSE) and higher correlation for latent flux yielding an average RMSE of 90 W m\(^{-2}\) and r exceeding 0.7, which largely outperforms METRIC and ESVEP by more than 11% and 30% for RMSE and Pearson correlation, respectively. Furthermore, the TSEB-PT has been constantly updated in order to improve the modelling scheme for thermal sharpening, and as reported in Guzinski et al. (2023) enhanced DMS-driven TSEB-PT at field scale achieved accuracy of 0.8 mm per day, which is our next-future goal to be implemented. Moreover, Sánchez et al. (2023) conducted extensive study on the performance of LST downscaling in Spain, and based on their validation results with in-situ measurements the DMS approach gave nearly two times smaller RMSE error compared to the 1-km S3 LST. In addition to the abovementioned literature review, in our co-authored paper we compared Sen-ET outcomes with other evapotranspiration products, including 3-km MSG SEVIRI and 70-m ECOSTRESS ET which on average gave less robust accuracy metrics than our 100-m retrievals (De Santis et al., 2022). These results and other authors’ findings moved us towards generation of 100-m ET dataset. Therefore, we have decided to apply the LST sharpening strategy in our ET workflow assuming its better performance in different land covers and climates compared to original 1-km S3-driven TSEB-PT. In the section 3.2 of the revised manuscript together with relevant research papers we provide more information on the performance of DMS procedure for estimating high spatial resolution evapotranspiration (lines 327-340; 358-363).

Considering the scarcity of eddy covariance towers over the Mediterranean catchments and time of interest (2017-2021) for our analysis, together with University of Ghent we have managed to gather in-situ observations at only eight eddy covariance sites (i.e., seven locations in Italy and one site in France) that provide long time-series of latent heat flux. In order to get more general conclusions on the results, we fully agree that the validation shall be performed including more in-situ EC towers represented by wider variety of landcover types, climate zones, and topography which is our future priority objective. This might be done by extending the spatial coverage of the ET data in order to increase number of available local flux measurements.


3) Introduction:

- Second paragraph of the introduction: when the authors review existing evapotranspiration models, they mention process-based (energy balance models) and data-driven (statistical models) approaches. The so-called contextual/semi-empirical approaches are missed. I recommend completing this state of the art by adding a few references to contextual methods.

Response 3a: Thank you for your comment. We have included the contextual methods for evapotranspiration retrieval. You can find our modifications in the second paragraph of the Introduction section (lines 46-49). In the revised version of the manuscript, we included some research papers on the contextual ET methods. They are as follows:


- Line 96: "many data-driven approaches have been proposed, relying on empirical relationships between 1-km surface temperatures and high-resolution explanatory variables derived from Synthetic Aperture Radar (SAR) and Visible Shortwave Infrared (VSWIR) sensors (Li et al. 2019; Mao et al., 2021; Pu and Bonafoni, 2023)." As none of the above references include SAR data I suggest this one: Amazirh et al.

Response 3b: Thank you very much for this suggestion. We have included this interesting research paper as a reference in the manuscript (line 102).

4) Eddy covariance measurements:

I could find no information on how the authors derived daily ET estimates from 30-min eddy covariance measurements. Equation (1) explains the unit conversion from W/m² to mm/day, but the aggregation of hourly eddy covariance measurements at the daily scale is not described at all (?).

Response 4: To be more precise with the description for daily ET aggregation procedure from 30-min ground observations at available EC towers, in the third paragraph of the section 2.2 (starting from line 184) we provide more detailed information on daily ET retrieval according to the strategy developed by the project partners from the University of Ghent working in hydrology domain. Additionally, we provide a reference paper of Hulsman et al. (2023) that explains the preprocessing procedure for the in-situ EC observations:


5) Spatio-temporal coverage of the dataset:

I am surprised by the relatively large and frequent gaps in the ET dataset due to cloud cover. I imagine that the S2 dataset composited over 10 days and the S3 dataset composited over 10 days separately have greater spatial coverage. I wonder if the relatively low spatio-temporal coverage of the ET dataset is associated with the temporal mismatch between S2 and S3 overpasses?

Response 5: As described in the Methodology section (lines 248-253; 306-309), we have generated daily ET maps with spatio-temporal coverage corresponding to daily Sentinel-3 (S3) LST data and 10-day Sentinel-2 composite product specially adjusted to S3 acquisition days. Indeed, in case of Sentinel-2 we minimized the cloud occurrence by means of temporal compositing, while S3 LST is more affected by overcast conditions. Sentinel-3 LST datasets were not composited and this is the main reason for relatively large spatiotemporal gaps in the daily ET product.

6) TSEB modeling

One of the difficulties in spatializing the TSEB over large areas is characterizing the aerodynamic resistance (linked to canopy height, leaf size, etc.) and the green component fg of the vegetation cover. Can you briefly present the range of values chosen for these key parameters for the main vegetation types in the basins studied?

Response 6: Thank you for this comment. Canopy aerodynamic resistance (Rx) and green vegetation cover (fg) are expressed as follows:

1. \( Rx = (C'/F) * \left[I_w/(u_{do}+z_{om})\right]^{1/2} \)
where $C'$ is derived from weighting a coefficient in the formulation for leaf boundary layer resistance over the height of canopy (it assumed to be equal to 90 s^{1/2} m^{-1}), F is local leaf area index (i.e., LAI/\textit{fc} with \textit{fc}: fractional vegetation cover), \textit{lw} is the effective leaf width, and $u_{d0}+z_{em}$ corresponds to the wind speed within the canopy-air interface (Norman et al., 1995).

2. $f_g=\text{FAPAR}/\text{FIPAR}$
where FAPAR is the fraction of absorbed photosynthetically active radiation obtained from the ESA Snap biophysical processor, and FIPAR corresponds to the fraction of photosynthetically active radiation intercepted by green and brown vegetation and it is expressed using following formula:

3. $\text{FIPAR} = 1 – \exp\left[-0.5*\text{PAI}/\cos\Theta\right]$ where $\text{PAI} = \text{LAI}/f_g$.

As shown above, the retrieval of $f_g$ is an iterative procedure that re-calculates FIPAR parameter until $f_g$ converges (Guzinski et al., 2020).

Green vegetation cover is driven by plant area index (PAI), FAPAR, and sun zenith angle ($\Theta$) derived from Sentinel-2 reflectance imagery. It means that $f_g$ values are estimated for each S2 pixel in space and time ranging from 0 to 1. Similarly to $f_g$ product, aerodynamic resistance at the canopy boundary layer is based on Sentinel-2 grid and changes in time since it is derived from Earth Observation inputs, such as LAI, ERA5 wind speed, and landcover information derived from ESA CCI LUT (e.g. \textit{lw}) following Guzinski et al. (2019).


7) Correction of input meteorological data for topography effects:

Line 337: "All extracted variables from the reanalysis dataset, except for wind speed, are corrected for terrain using the SRTM DEM product"

Line 440: "The distribution of solar radiation, wind speed, and air temperature gradients are less influenced by a landscape complexity over mountain plateau than over steep slopes, and thus coarse resolution ERA5 might be more representative for … "

Line 474: "The ET models are controlled by climate inputs derived from 31-km fields…"

The above statements seem contradictory. Can you please describe how solar radiation and air temperature are downscaled at 100 m resolution using the DEM? Both variables have a very strong effect especially in areas of complex topography such as the basins studied?

**Response 7:** Indeed, air temperature and solar radiation from ERA5 data were enhanced with SRTM DEM. While shortwave radiation was corrected for illumination conditions using elevation, air temperature (TA) originally derived at 2-m height was recalculated to the blending height of 100 m using the elevation product and standard lapse rate of 6.5 K/1000 m. The blending height for low-spatial resolution air temperature is assumed to be more representative rather than TA at the height of 2 m due to weaker impact of land-atmosphere interactions at 100-m above ground surface. We expect a strong impact of those variables over
complex areas, such as mountains where most EC towers are located in this study. To be more consistent, we have added some text (lines 234-235; 365-370) to explain the utility of DEM in the ERA5 input data adjustments together with a relevant research paper:


8) Shadows effects

Line 450: «The poor accuracy at forested sites might be related to their complex tree structures and multilayer composition which is not considered in Sen-ET». Since the authors are evaluating their product at Puechabon site, I suggest referring to Penot et al. (2023). Estimating the water deficit index of a Mediterranean holm oak forest from Landsat optical/thermal data: a phenomenological correction for trees casting shadow effects. IEEE Journal of Selected Topics in Applied Earth Observations and Remote Sensing.

Response 8: Thank you very much for this suggestion. We included this publication in our manuscript (line 485).

9) Discussion of sources of uncertainty in the ET dataset (Lines 465-473)

Another source of uncertainty that should be mentioned and discussed is the intrinsic limitation of downscaling methods of LST data using reflectances as high resolution ancillary information.

Response 9: We fully agree with the reviewer’s comment. We have included this aspect in the indicated paragraph of the revised manuscript (lines 510-514). We also added some reference papers to explain limitations of reflectance bands as predictors for thermal downscaling as depicted below:


Edits:

- Once defined, acronyms must be used systematically (this problem appears in many places in the text). Also acronyms should be defined only once.
- There is an acronym for Languedoc Roussillon but not for the study basin (Herault basin)?
- Composted/composting (line 17, fin 2, 396, line 486)
- Unit is missing for RMSEs at line 584

Response to edits: Thank you. We have corrected the manuscript considering the abovementioned issues. In addition, we updated the table with a list of abbreviations and acronyms (page 27-28).