The First Hillslope Thermokarst Inventory for the Permafrost Region of the Qilian Mountains

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Abstract:

Climate warming and anthropogenic disturbances result in permafrost degradation in cold regions, including in the Qilian Mountains. These changes lead to extensive hillslope thermokarst (HT) formation, such as retrogressive thaw slumps, active-layer detachment slides, and thermal erosion gullies. These in turn cause, e.g., degradation of local vegetation, economic losses, infrastructure damages, and threats to human safety. However, despite its importance, there is currently no thermokarst inventory for the Qilian Mountains. Through manual visual interpretation and field validation, we therefore produce the first quantification of HT features. We count a total of 1064 HT features, with 67% located in the upper reaches of the Heihe River Basin, which encompasses ~13% of the Qilian Mountains region. We furthermore document that 82% of the HT was initiated in the last 10 years. The thermokarst terrain is observed primarily in areas with shallow active layer depth (average thickness: 2.98 m), on northern shaded slopes of 3–25°, with low solar radiation and moderate elevations ranging from 3200 to 4000 m. This first inventory of HT features is an important and missing piece in documenting changes on the Qinghai-Tibetan Plateau, and this new dataset also provides an important basis for further studies on, e.g., quantitative assessment losses caused by HT. The datasets are available from the National Tibetan Plateau/Third Pole Environment Data Center and can be downloaded from https://doi.org/10.11888/Cryos.tpdc.300805 (Peng and Yang, 2023).

1 Introduction

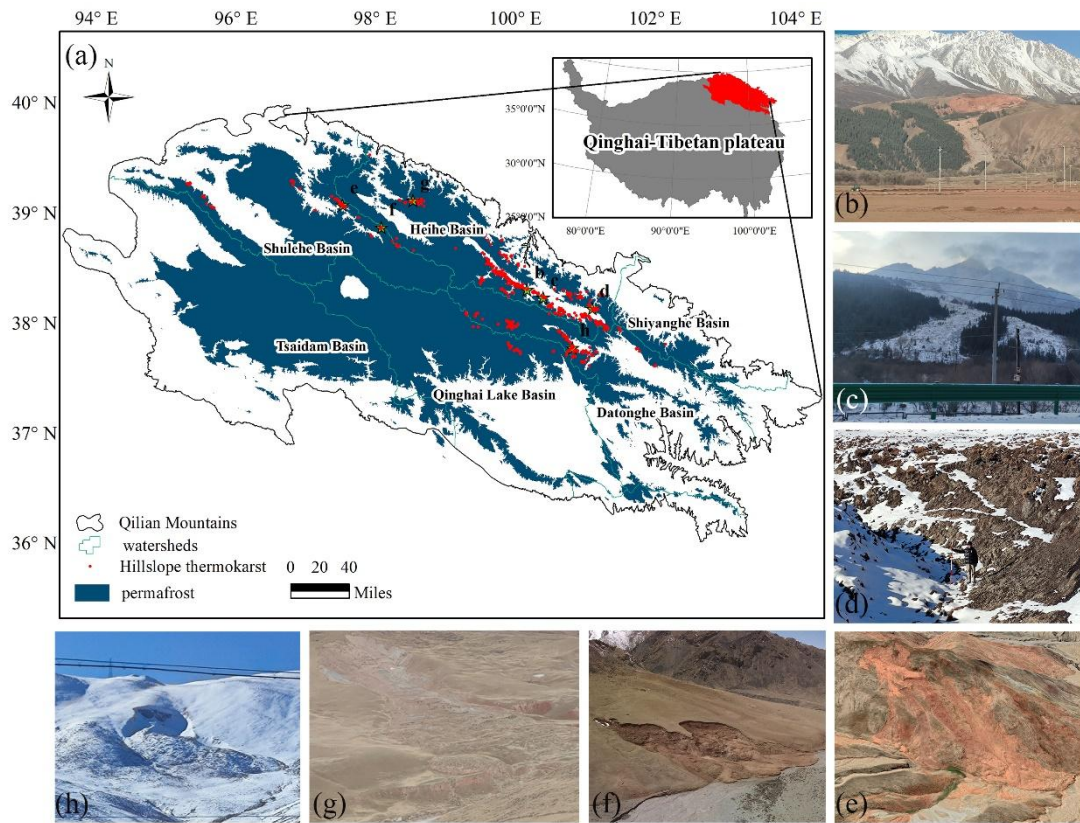
The Qilian Mountains are located in the northeastern part of the Qinghai-Tibetan Plateau, at the confluence of three major geographical regions that include the eastern monsoon zone, the northwestern arid zone, and the alpine zone of the Qinghai-Tibetan Plateau. The Qilian Mountains play an important role in maintaining the ecological balance of the Tibetan Plateau, stopping the southward progression of deserts, and maintaining the stability of the oases in the Hexi Corridor. Due to its unique geographical and environmental characteristics, permafrost is widespread and underlies about 50% of the area (Ran et al., 2021). Permafrost has an important role in storing frozen water, thereby contributing to water conservation (Wang et al., 2022). These roles can aid in inland river runoff recharge, which is crucial to regional ecology, production, and life. Due to climate warming and human activities, significant permafrost degradation results in the frequent occurrence of thermokarst, representing a serious threat to ecological security and adversely impacts the environment and human beings (Li et al., 2022a).

Despite the importance of thermokarst processes and their potential geohazards, the distribution of thermokarst landscapes is currently mostly undocumented. The available distribution of thermokarst in the Northern Hemisphere, including retrogressive thaw slumps (RTSs), thermokarst lakes, and other terrain features, represents mainly probabilistic estimates (Olefeldt et al., 2016; Yin et al., 2021; Huang et al., 2023; Yang et al., 2023). Muster et al. (2017) determined the distribution of circumpolar Arctic permafrost lakes and ponds from 2002–2013 at a resolution of 5 m using optical remote sensing, satellite (Geo-Eye, QuickBird, WorldView-1 and -2, KOMPSAT-2), and radar imagery (TerraSAR-X), but temporal inconsistencies make comparisons in time and space difficult. At the regional scale, the techniques and spatial resolution of remote sensing imagery currently used at different study areas are inconsistent, e.g., estimating the distribution and development of RTSs on Banks Island, Canada, based on the interpretation of a Google Earth Engine Timelapse dataset (Lewkowicz and Way, 2019). Satellite imagery at 30-m resolution from Landsat has been used to analyze RTSs and thermokarst lakes in the region of Alaska within the circumpolar Arctic, eastern Canada, and Siberia (Nitze et al., 2018). Nicu et al. (2023) employed orthorectified imagery with a resolution of 5 meters to visually interpret and identify cryospheric hazards (such as thaw slumps and thermo-erosion gullies) in Nordenskiöld Land, covering an approximate area of 4000 square kilometers in the Svalbard Archipelago. The permafrost zone of the Qinghai-Tibetan Plateau (QTP) has been a site of thermokarst geomorphology research in recent years. For example, combining field surveys and SPOT-5 satellite data for August 2010, a total of 2,163 thermokarst lakes and ponds were recorded within 10 km on either side of the Chumar River to Fenghuo Mountain of the Qinghai-Tibet Railway, with a total surface area of 1.09×10^7 m^2 and ranging in size from 100 m^2 to 4.49×10^5 m^2 (Luo et al., 2015; Niu et al., 2014). In the Beiluhe region of the central QTP, the number of RTSs increased from 124 to 438 between 2008 and 2017, with an approximate 9-fold increase in area (Huang et al., 2020; Luo et al., 2019). The latest results show that the number of RTSs on the QTP is 2669, but for the Qilian Mountains in the northeastern part of the region, only 6 (Luo et al., 2022) or as many as 15 are documented (Mu et al., 2020). A lack of a thermokarst inventory in this region is therefore evident, representing a crucial gap in the RTSs inventory on the QTP.

Frequent occurrence of hillslope thermokarst hazards due to permafrost degradation,

with significant ecological impacts on the Qilian Mountains. The ecological environment of the permafrost areas has a significant impact, and there is a direct correlation between human activities and major permafrost engineering problems, including uneven subsidence of infrastructure, slumps, and cracks. Meanwhile, there is little to no information regarding hillslope thermokarst (HT) features such as RTSs, active-layer detachment slides, and thermo-erosion gullies (Gooseff et al., 2009) in the Qilian Mountains. HT refers to a specific type of thermokarst formation that occurs in permafrost regions. While it is similar to regular thermokarst features, what distinguishes hillslope thermokarst is its occurrence on sloped terrain or hillsides, where permafrost thaw leads to slope instability. This can result in various landforms like retrogressive thaw sumps, thermo-erosion gullies, or active layer detachments, affecting the stability and shape of hillslopes in permafrost regions. These features can significantly impact the landscape and have implications for infrastructure, ecosystems, and land use in areas affected by hillslope thermokarst processes (Kokelj and Jorgenson, 2013; Olefeldt et al., 2016; Gooseff et al., 2009). Thus, the urgent need to survey and quantify these undocumented HTs in the Qilian Mountains motivates and represents the goal of this study.

2 Study Area

The Qilian Mountains are located at the northern edge of the QTP, with an average elevation of 3855 m. The region is underlain by permafrost and seasonally frozen ground (36–40°N and 94–104°E, Figure 1a), with a permafrost area of 94,235 km^2 that accounts for 49% of the study domain. Characterized by both an alpine mountain climate and a temperate continental monsoon climate, the mean annual air temperature is 0.30°C (Jin et al., 2022) with high precipitation variability and higher amounts in the southeast during the thawing season of June to September (Chen et al., 2013; Li et al., 2022b). Due to human activities, climate change, and earthquakes, permafrost instability in Qilian Mountains has gradually increased, resulting in HT formation including RTSs, active-layer detachment slides, and thermal erosion gullies, which pose a direct threat to the local economy, ecology, and safety.

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119 **Figure 1** The location of the study area and a) its HT distribution (Qilian Mountains
120 permafrost extent data is from Sheng et al., 2020), and b)–h) HT features obtained
121 from different watersheds during our field surveys with the exception of e) a Google
122 Earth image, as this site is too difficult to access, the positions corresponding to b)-h)
123 have been labeled in a).

124

3 Data Sources

126 We collected and collated validated satellite imagery available starting in 1999 for
127 temporal detection of the onset of the HT formation. These data include unmanned
128 aerial vehicle imagery (e.g., Figure 1b–h) and 30 m resolution digital elevation model
129 data from the Shuttle Radar Topography Mission (Farr et al., 2007). A combination of
130 Omap and Google Earth software was used to detect the location of HT occurrence,
131 and Wayback imagery provided by ESRI was used to access high-resolution (<1 m)
132 satellite imagery and Jilin-1 satellite imagery (0.75 m) provided free of charge by
133 China Commercial Satellite Corporation to aid in the identification (Table 1). In
134 addition, we used digital elevation model data to calculate variables such as slope and
135 topographic position index (TPI) of the HT. The TPI is calculated as follows (YU
136 Hong, 2001):

$$\text{TPI} = \log_{10}\left(\frac{E}{Mean\,E} + 1\right) \times \left(\frac{S}{Mean\,S} + 1\right) \quad (1)$$

138 where E is the elevation (m), S is the slope (°), and *Mean* indicates that the mean
139 value for the entire Qilian Mountain region is used.

140 To further analyze the distribution of HT and the analogous environmental variables,
141 We obtain seismic data from the U.S. Geological Survey
142 (https://earthquake.usgs.gov/earthquakes/search/) describing earthquakes, including

their timing, epicenter location, and magnitude. To categorize vegetation types into deciduous-coniferous forests (DCF), undergrowth (U), alpine scrub meadow (ASM), alpine meadow (AM), alpine grassland (AG), alpine vegetation (AV), and non-vegetated area (NA), based on data from the Resource and Environment Science and Data Center (https://www.resdc.cn/data.aspx?DATAID=122). To assess the relationship of air temperature and precipitation with HT, we download monthly mean air temperature and precipitation at 2 m above ground level from the fifth generation of the European Centre for Medium-Range Weather Forecasts (ECMWF) reanalysis (ERA5; https://cds.climate.copernicus.eu/cdsapp#!/dataset/reanalysis-era5-land-monthly-means?tab=overview).

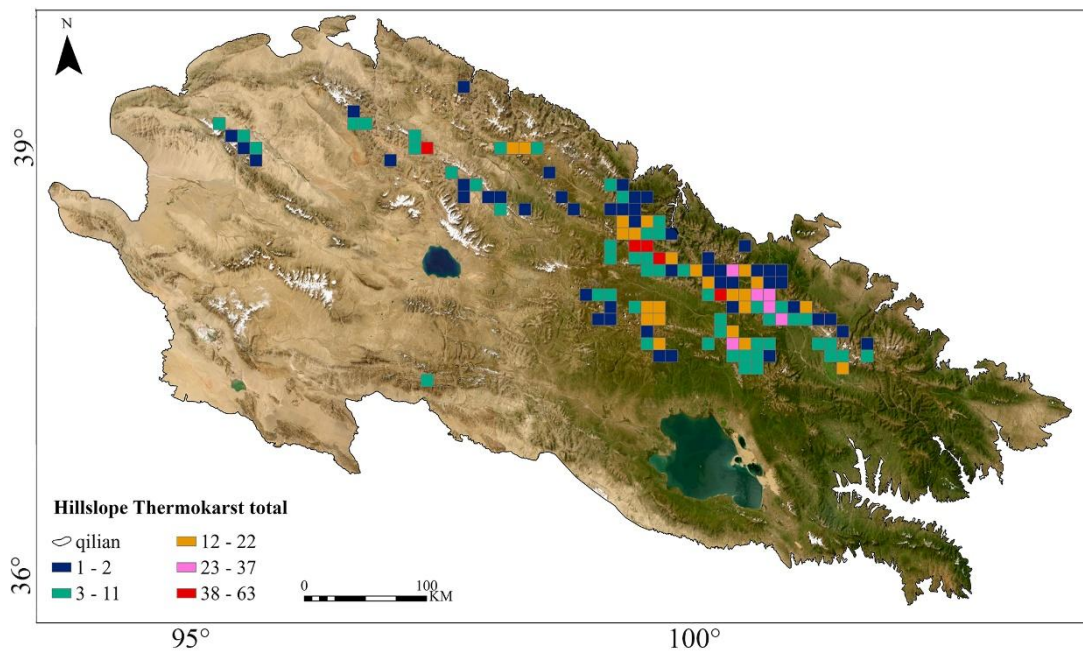



Figure 2 Frequency distribution of HT on the Qilian Mountains. The concentration of HTs is shown per 100 km^2 grid cell.

4 Methods

4.1 Manual Mapping

We first quantified and mapped HT via remote sensing observations. Most occurrences of HT in the permafrost region of the Qilian Mountains since 2000 were compiled by visual interpretation in Google Earth Pro and Omap. They were also aided by high resolution (<1 m) observations from Esri Wayback Imagery, which archives all published versions of world imagery (Table 1). We used a fishnet with a mesh size of 1×1 km to segment the latest satellite imagery for the entire Qilian Mountains to quantify HT mesh by mesh. RTSs are often horseshoe shaped, tongue shaped, elongated, branched, and circle chair-shaped (Lantuit and Pollard, 2008; Yin et al., 2021), characterized by a steep backwall, low-angle bottom, and a tongue of displaced saturated soil (Lantz and Kokelj, 2008; Nicu et al., 2021). These features are tonally and morphologically different from their surroundings in color satellite images during the thawing season. RTSs also produce folded textures due to soil

accumulation, which appear as laterally folded stripes on imagery. Active-layer detachment slides are a common shallow landslide in permafrost areas. Their morphological characteristics vary based on vegetation cover, slope, and permafrost conditions, but common features include highly disturbed slopes, lateral shear zones, and fracture zones formed after the sliding of the active layer. (Lewkowicz, 2007). We detected and sketched these features based on morphological, tonal, textural, shading, and other characteristics on remote sensing images, and then digitized their morphological features into polygonal data. Although the accuracy of this type of visual interpretation is relatively high, some HT features can be missed via this manual interpretation. To reduce such errors, satellite images of the similar period from different sources were evaluated four times using the same methods to ensure accurate results. The date of the satellite image when perturbations caused by HT can be first observed was defined as the initiation year of a particular HT feature. Depending on the initiation year, HT is categorized as occurring before 2010, 2010–2015, or after 2015. To observe the temporal evolution of HT features, we used the initiation year and retraced historical images covering the Qilian Mountains, a process that also helped us distinguish between HT features and one-time transient landslides.

4.2 Field Verification

Similar HT can have different morphological characteristics due to different triggers. It is thus difficult to identify the type of HT simply through imagery. In addition, after an initial trigger and HT formation, thermokarst can evolve into different types. For example, active layer detachment slides may transition into RTS due to the exposure of subsurface ice at the trailing edge and water erosion due to thawing, which can cause the RTS to further progress into mudflows. Therefore, with visual interpretation based on imagery only providing individual snapshots, it is essential to also conduct field surveys as a validation exercise. We conducted a total of three field surveys in winter 2022, and spring and summer of 2023. Field work covered the Shiyanghe basin, Heihe basin, Datonghe basin, Qinghai Lake basin, and Shulehe basin. Due to the harsh climatic conditions and accessibility issues in the Qilian Mountains, unmanned aerial vehicles were used to survey and verify hard-to-reach areas.

Table 1 List of the data used for manual interpretation and mapping for HTs.

Software Platform	Time Span	Resolution	Data Sources
Google Earth pro	1999-2022	0.6-15 m	Quickbird, IKONOS, etc.
Omap	since 2021	<1 m	GF-2
ESRI World Imagery	since 2014	<1 m	WordView-3, WordView-4, etc.
Jilin-1 website	2022	0.75 m	Jilin-1
UAV images	Feb., Apr., May 2023	~15 cm	Field Surveys

4.3 Morphological and Spatial Statistical Analysis

A landscape shape index (LSI) can be quantified to characterize shape complexity by calculating the degree of deviation of a given patch from a circle or square of the

same area. To quantify the shape characteristics of HT features, two LSIs are calculated as follows:

$$LSI_{square} = \frac{0.25P}{\sqrt{A}} \quad (2)$$

$$LSI_{round} = \frac{P}{2\sqrt{\pi A}} \quad (3)$$

where P is the perimeter (m) and A is the area (m^2). The closer the values of LSI$_{square}$ or LSI$_{round}$ are to 1, the more square or round the shape of the HT feature is, respectively.

To further investigate the spatial distribution of HT, we computed a global Moran's index, z-score and p-value to determine whether there is autocorrelation in the spatial distribution of HT. Where p-value and z-score are used to measure statistical significance, when p-value < 0.01 and z-score > 2.58, it means that there is a 99% probability that HTs are clustered within the study area, and the smaller the p-value and the larger the z-score, the greater the probability that such spatial patterns are clustered. Moran's index ranges from -1 to 1, with negative values meaning negative correlation, positive values meaning positive correlation, and 0 denotes that the spatial objects in the study area are independent of each other. Additionally, the closer the index is to 1, the more clustered the HT features are, and the closer of the index is to -1, the more dispersed the HT features are. To delineate the regions that may have spatial autocorrelation (Bivand and Wong, 2018), we further process local autocorrelation on this basis. The local autocorrelation regions are divided into four types: The local autocorrelation analysis categorizes regions into four types based on the local Moran's index: High-High (HH) clustering, High-Low (HL) clustering, Low-High (LH) clustering, and Low-Low (LL) clustering. HH signifies a region with both a higher amount of HT and neighboring regions also having a higher amount of HT; HL indicates a region with a higher amount of HT surrounded by neighboring regions with a lower amount of HT; LH indicates a region with a lower amount of HT neighboring areas with a higher amount of HT; and LL represents a region with both a lower amount of HT and neighboring regions with a lower amount of HT. Although the methods described above can identify global and local spatial autocorrelation, respectively, they are unable to identify clusters of concentrated HT features. We therefore also apply hot spot analysis, which is another effective way of exploring the characteristics of local spatial distributions. All the above techniques are based on spatial statistical analysis functions of ArcGIS.

To explore the effects of climate on HT, we obtained the monthly mean air temperature and precipitation at 2 meters above ground level from ERA5 over the period 2000–2020 and calculate their annual spatial means and standard deviations (Figure 6).

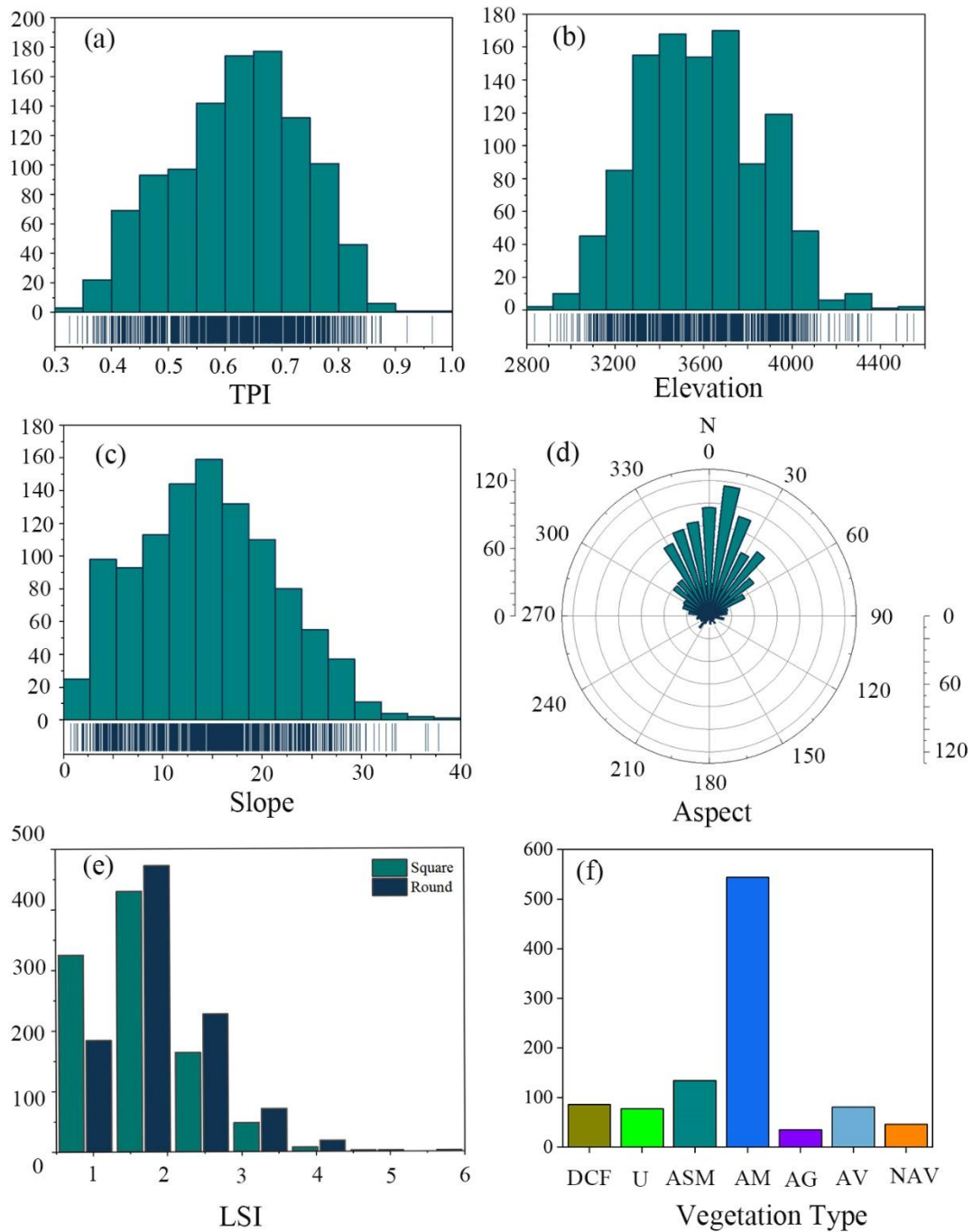
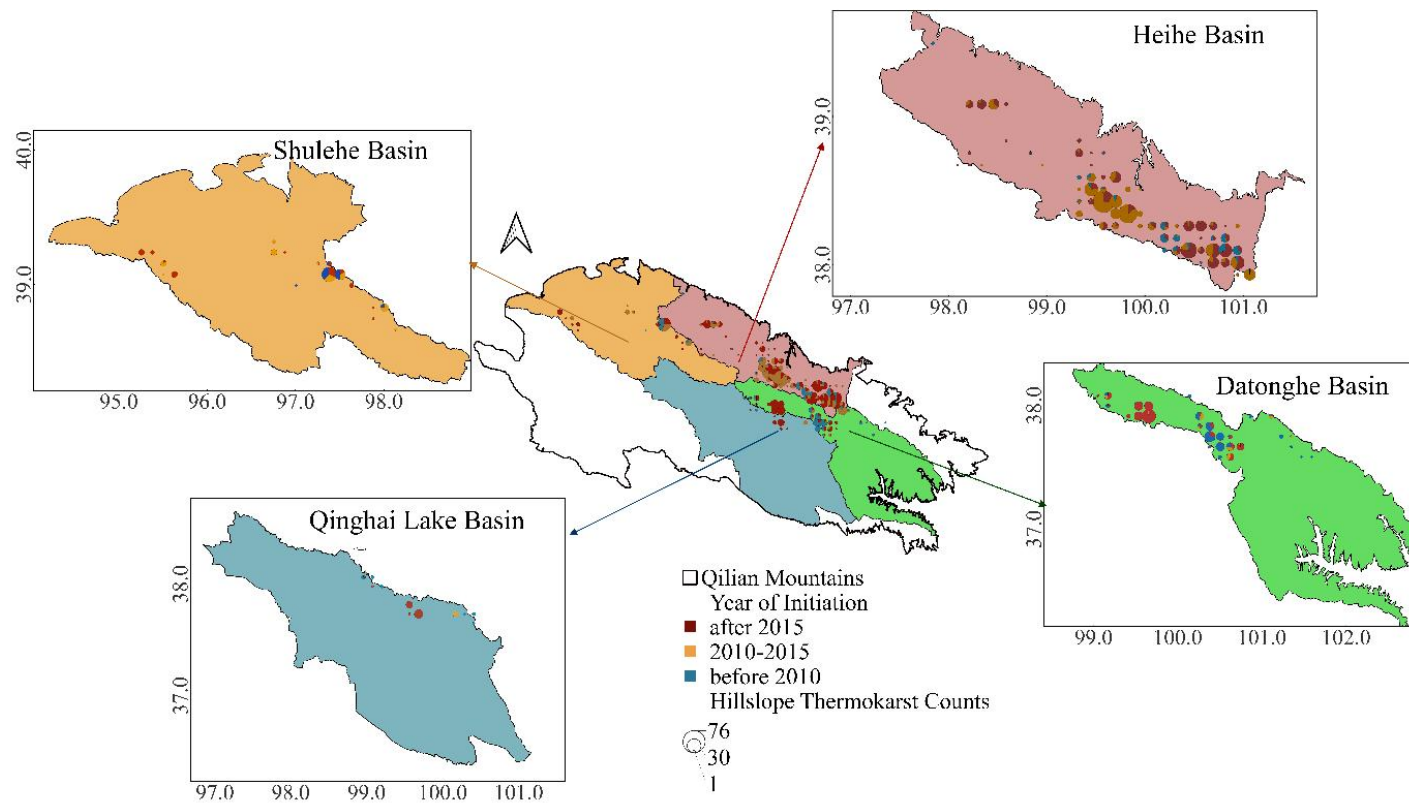
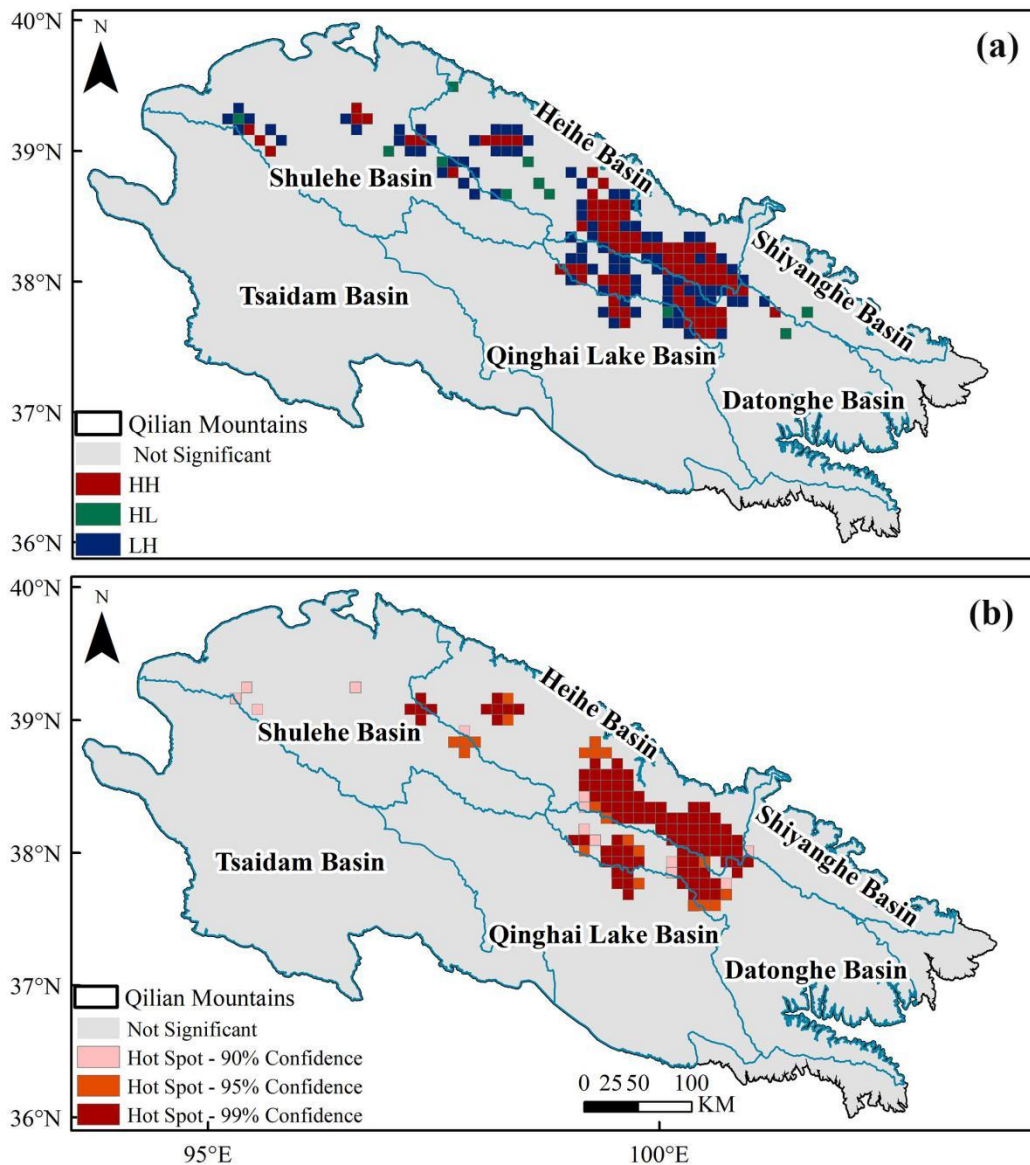



Figure 3 The number of HT terrain features (y-axes) of HT as categorized by (a) topographic position index (TPI), (b) elevation, (c) slope , (d) aspect, (e) landscape shape index (LSI), and (f) vegetation type including deciduous coniferous forests (DCF), undergrowth (U), alpine scrub meadow (ASM), alpine meadow (AM), alpine grassland (AG), alpine vegetation (AV), and non-vegetated area (NA); the blue vertical lines at the bottom of panels a–c represent the number of HT features in each x-axis bin.

5 Results

259 Our inventory of HT includes the Heihe Basin, Shulehe Basin, Datonghe Basin,
260 Shiyanghe Basin, Qinghai Lake Basin, and Tsaidam Basin within the Qilian
261 Mountains, with a total of 1064 HT features. In any 100 km^2 grid cell, the maximum
262 density of HT is 63 (Figure 2).This density is lower than the 68 per 25 km^2 in the
263 central Tibetan Plateau reported by (Luo et al., 2022) and 88 per 25 square km^2 on
264 Banks Island, Canada from (Lewkowicz and Way, 2019). 67% of these HT features
265 were identified in the Heihe River basin, followed by the Datonghe River Basin,
266 accounting for 19%. The HT distribution in these river basins is irregular,
267 corroborated by a positive statistically significant Moran's index value of 0.3, p-value
268 of 0.00001 and z-score of 32.5. Of all the HT features, the largest is 58 ha, the
269 smallest area is 0.01 ha, with most being smaller than 10 ha. The average area is 1.75
270 ha, with a total area of 1708 ha.

271 The spatial distribution of RTS on the QTP is strongly controlled by terrain factors
272 such as the elevation, slope, TPI, and aspect (Luo et al., 2022). The statistical results
273 indicate that HT is observed at elevations ranging from 2,835 to 4,550 m. However,
274 90% of HTs are more likely to occur at elevations ranging from 3,200 to 4,000 m in
275 the middle/high elevation area of the Qilian Mountains. HT tends to occur on
276 north-facing slopes (Figure 3b and 3d), with slopes ranging from 3° to 25° (Figure 3c).
277 In addition, the TPI shows that ~85% of the HTs occur mainly between 0.5 and 0.8
278 (Figure 3a), suggesting that they commonly occur in locations that are lower than
279 their surroundings. Both LSI indices suggest that 75% of HT has values close to 1.0
280 (Figure 3e), indicating that most HT is simple in shape and compact in morphology
281 (Niu et al., 2016). Alpine meadow areas contain ~53% of HT, followed by alpine
282 scrub meadows, which contain 13% (Figure 3f).

283 The initiation years of HT features are variable across the study area. 187 HT features
284 (18%) were identified before 2010, and the remaining 82% in the last 10 years. 392
285 sites (37%) were initiated in 2010-2015 and 482 (45%) after 2015. Much of the newly
286 initiated HT occurred in the Heihe basin and the middle and upper reaches of the
287 Datonghe basin (Figure 4), which is also a HT hotspot region. The recent increase in
288 HT can be attributed to the anomalous weather conditions in the corresponding years.
289 The association between newly observed HT and meteorological data indicates a
290 sudden HT increase in years with unusually high temperatures during the thawing
291 season (Figure 6).

292
293 **Figure 4** The timing of HT initiation within 100 km^2 grid cells.

294
295 **Figure 5** (a) Spatial auto-correlation indicating high-high (HH), high-low (HL), and
296 low-high (LH) clustering, and (b) hotspot analysis where the different colors represent
297 the confidence levels.
298

6 Discussion

6.1 Drivers of HT in the Qilian Mountainous

6.1.1 Permafrost Conditions

Formation of HT is facilitated by thick subsurface ice and various internal and
external environmental conditions (Stephani et al., 2023). Permafrost stability in
~80% of the permafrost area of the Qilian Mountains is predominantly transitional,
and higher permafrost temperatures (Ran et al., 2021) will exacerbate the climate
sensitivity of this area (Lewkowicz and Way, 2019; Patton et al., 2021) leading to
melting of the subsurface ice and an increase in active layer thickness, thus decreasing
the stability of the slope (Behnia and Blais-Stevens, 2018). This is also supported by
our finding that ~90% of HT occurs in the transition zone between permafrost and

310 seasonally frozen soil where mean annual ground temperature is greater than −1°C.
311
312 *6.1.2 Environmental Factors*

313 Topographic conditions facilitate the formation of subsurface ice and the continuous
314 development of HT. At elevations below 5100 m on the QTP, aspect dominates the
315 distribution of permafrost. More permafrost underlies regions of shaded, north-facing
316 slopes than sunny south-facing slopes (Ran et al., 2021). Indeed, we find that ~95% of
317 Qilian Mountain HT is found on north-facing slopes where it also enhances vegetation
318 growth and soil moisture storage (Jin et al., 2009). Lower solar radiation, higher
319 permafrost ice content, and shallow active layer thickness (Lacelle et al., 2015; Ward
320 Jones et al., 2019) also enables HT formation (Luo et al., 2022; Niu et al., 2016; Xia
321 et al., 2022). We find more than half of the HT occurs in alpine meadows, which
322 require more water content than alpine steppes (Yin et al., 2017) and consequently
323 also results in more ground ice development under this vegetation type. We
324 determined that ~90% of HT in the Qilian Mountains occurs on 3° to 25° slopes. Low
325 and gentle slopes are favorable for groundwater pooling (Luo et al., 2022), whereas
326 slopes greater than 16° are relatively steep and therefore not conducive to
327 groundwater enrichment for ice formation, but such slopes also provide dynamic
328 conditions for active layer detachments and collapsing ground (Wang, 1990). We also
329 observe more HT initiation at locations that are lower compared to their surroundings,
330 as such depressions favor the accumulation of snow and rainwater (Stieglitz et al.,
331 2003) and prevent heat loss from the soil. This encourages melting of subsurface ice
332 (Zhang, 2005) at the base of the active layer and, after an unstable layer is formed
333 between the permafrost and the active layer, the overlying soil can slide along the
334 slope (Patton et al., 2021).

335 The Qilian Mountains were and are still formed by the ongoing collision of the Indian
336 Ocean Plate and the Eurasian Plate, resulting in the Qilian Mountains-Hexi Corridor
337 active fault system (Xiong et al., 2017) that has seen nearly 400 earthquakes of
338 magnitude 2 or greater over the past two decades. In particular, the high seismic
339 activity of the Heihe, Shiyanghe, and Datonghe Basins (Figure 7a) represents a
340 potential threat to the safety and integrity of current and future infrastructure in the
341 region. During our field investigations we found a nearly 3 km long and 2 m deep
342 slope fracture caused by a 6.9-magnitude earthquake in 2022, resulting in a massive
343 exposure of subsurface ice and the collapse of the Lanzhou-Xinjiang High Speed Rail
344 Tunnel (Figure 7b and c). The occurrence of an earthquake can result in an
345 instantaneous increase in pore water pressure and sliding forces that reduce slope
346 stability and potentially leads to a massive exposure of subsurface ice (Niu et al., 2016;
347 Xia et al., 2022), sediment liquefaction (Dadfar et al., 2017), and permafrost warming
348 due to the seismic vibrations. These vibrations lead to cracking and deformation of the
349 ice layers within the permafrost, releasing moisture and heat, consequently resulting
350 in a temperature rise of the permafrost. Additionally, earthquakes can induce the flow
351 of pore water within the permafrost, further influencing its temperature (Che et al.,
352 2014), creating the ideal setting for active-layer detachment slides. The clustering of
353 earthquake activity in the lower left corner of Figure 7a is not associated with HT
354 features, because there is no permafrost in this region.

355
356 *6.1.3 Climate Factors*

357 Extreme summer temperatures and precipitation have been identified as triggers for
358 the initiation of RTSs in many Arctic permafrost zones (Balser et al., 2014; Kokelj et

al., 2015; Lewkowicz and Way, 2019; Segal et al., 2016). Given our finding that 82% of HT was initiated in the last decade (Figure 4), mostly during 2010-2015 and after 2015, we used ERA5 to determine the temperature and precipitation characteristics for the Qilian Mountains over the last 20 years (Li et al., 2022b) (Figure 6, the square symbols). The mean thawing season air temperatures in 2010 and 2016 were higher than in other years (Figure 6a, red square symbols). A warming thaw season could lead to thaw consolidation at the base of the active layer or to higher porewater pressure in the transient thaw layer, reducing the effective shear strength, and causing slope failure (Lewkowicz and Way, 2019). The anomalous air temperatures during the thawing season could accelerate permafrost thaw and expose ice-rich permafrost, thus leading to new HT (Figure 6a, dark brown and dark red bars, respectively). Rainfall infiltration may transfer heat to the top layer of permafrost and induce melting of ground ice in ice-rich transient layers, which would increase the porewater pressure at the active layer-permafrost interface and thereby trigger formation of HT (Luo et al., 2022). However, precipitation variability during the thawing season does not match HT formation (Figure 6b). Despite high precipitation in both 2007 and 2018 (Figure 6b, blue squares), no initiation of HT was found to subsequently coincide with these peaks (blue squares), and precipitation also does not explain the significant initiation of HTs between 2010-2015, nor after 2015 (Figure 6b, yellow and red bars, respectively). The same conclusion also applies to the other three sub-regions—Hoh Xil Mountain, Maqu county, and Honglianghe—and it could be speculated that the nature of the soils on the QTP may instead play a role (Luo et al., 2022).

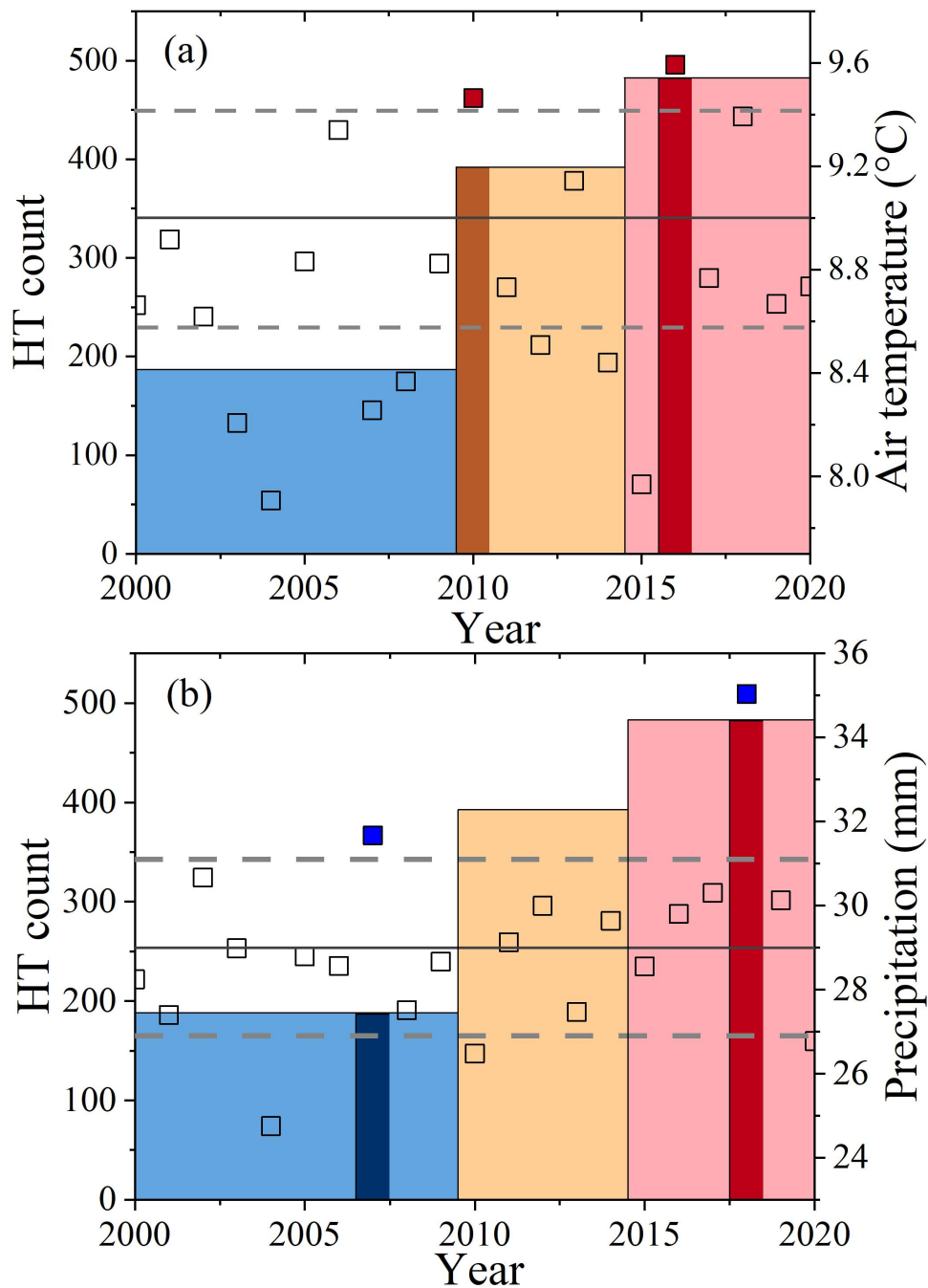



Figure 6 The relationship between HT numbers (unequal width bars, the darker colors represent years with extreme weather events.) and (a) temperature and (b) precipitation in the thawing season from 2000 to 2020 (square symbols, the red squares and the blue squares represent the extreme weather events.). The solid horizontal line represents the mean air temperature and precipitation, respectively, and the dashed line represents ±1 standard deviation.

6.1.4 Human Activities

Extensive and increasing human activities have been shown to significantly accelerate permafrost degradation (Cheng & Jin, 2013; IPCC, 2019). The total population of the Qilian Mountains is about 60,000, there are approximately 1,000 metal, energy, and other types of mineral deposits (National Mineral Properties Database 2021 Edition,

http://data.ngac.org.cn/mineralresource/index.html), and there are ~8,000 km of railroads and highways. The core of this human activity is concentrated on the eastern side of the Qilian Mountains, which generally coincides with the spatial distribution of the HT hotspots we documented. This qualitatively suggests a role of human activities on HT from, e.g., engineering disturbances, vegetation degradation due to overgrazing, etc. (Sharkhuu et al., 2007). Establishing the impact of human activities on HT quantitatively is still a difficult challenge, but our identification of the location and timing of HT formation is a first, important step for further future studies, especially on the socioeconomic development in the region.

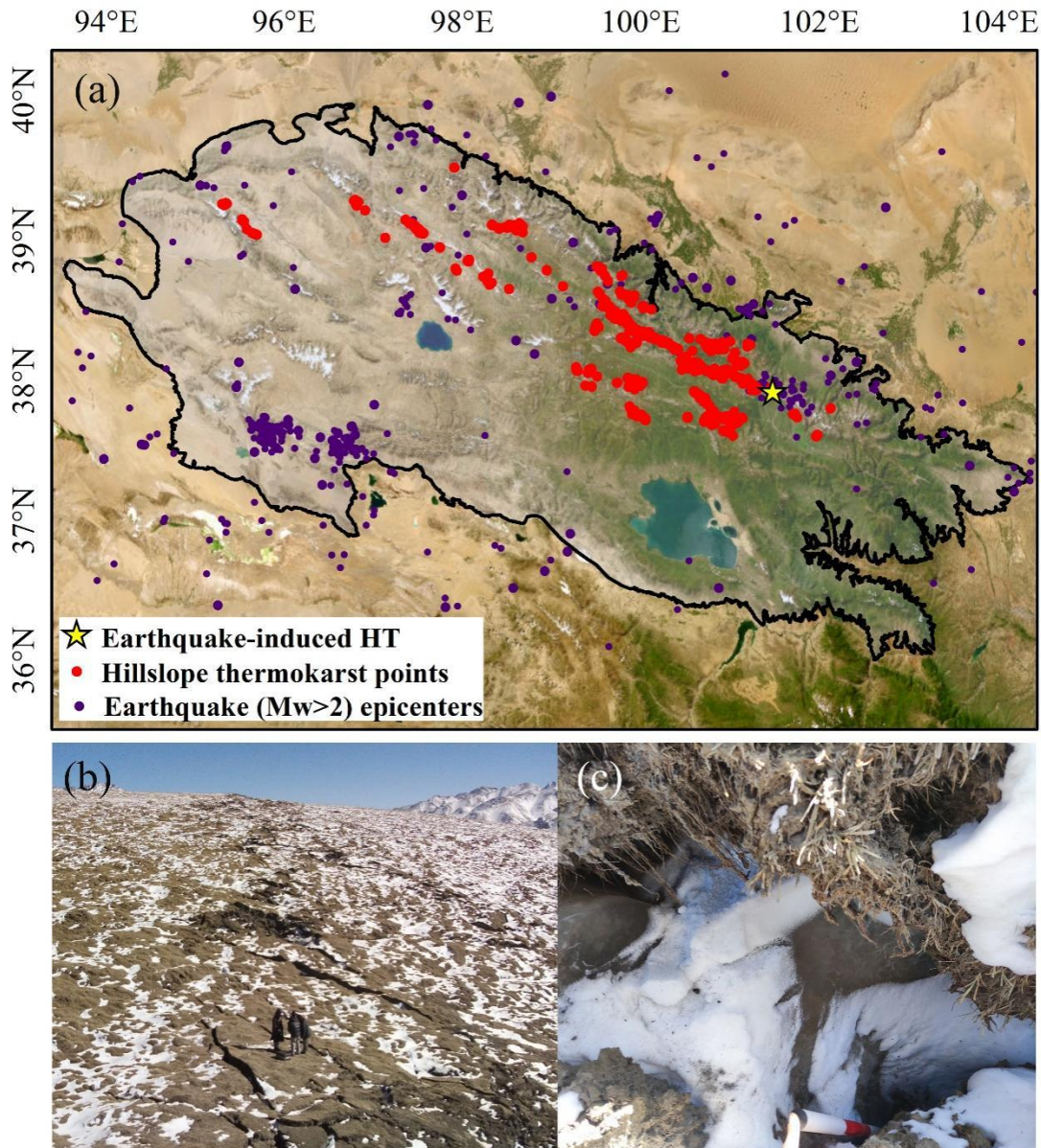



Figure 7 Qilian Mountains showing (a) the location of HT locations and earthquakes with magnitude >2 in the last 20 years, (b) slope fractures caused by earthquakes, and (c) exposed subsurface ice.

7 Data Availability

DEM data can be accessed through NASA

(https://www.earthdata.nasa.gov/sensors/srtm). Landsat5-8 data are available from USGS (https://www.usgs.gov/) and Sentinel-2 from ESA (https://www.esa.int/), and can be downloaded through Google Earth Engine. Esri World Imagery can be accessed through Esri Wayback Imagery at: https://livingatlas.arcgis.com/wayback. Some GF-2 imagery is also available online through Omap software (https://www.ovital.com/283-2/), and high resolution 3D satellite imagery of the Qilian Mountain region can be viewed in Google Earth software. High-resolution satellite images captured by the Jilin-1 satellite in China can be viewed by accessing https://www.jl1mall.com/rskit/. The HT inventory for the Qilian Mountains can be freely downloaded from the National Tibetan Plateau/Third Pole Environment Data Center (https://doi.org/10.11888/Cryos.tpdc.300805, Peng and Yang, 2023).

8 Conclusion

This study used visual interpretation and field investigations for repeated verification to investigate HT in the permafrost zone of the Qilian Mountains. We successfully produced the first HT inventory for this area, and found it contains a total of 1064 HT features. The area of these features ranged from 0.01 to 58 ha, with an average of 1.75 ha. Thermokarst is primarily concentrated at the junction between the upstream portion of the Heihe River Basin and the mid and upstream portion of the Datonghe Basin. Within a 10×10 km area, thermokarst has a maximum density of 63 features, affecting an area of ~20 km^2. HT in the Qilian Mountains is more likely to occur on north-facing shaded slopes, at elevations between 3200–4000 m, slopes of 3–25°, 0.5<TPI<0.8, and in alpine meadow vegetation. Based on long-term satellite imagery, 874 new HT features were initiated after 2010, accounting for 82% of the total HT. Recent higher air temperatures during the thawing season are likely important reasons for the intensification of HT formation in the Qilian Mountains, while precipitation does not seem to play a role. This first HT inventory for the Qilian Mountains will be fundamental for quantitative assessments that explore the exact causes and underlying thermokarst processes, and ultimately allow for better identification prediction of areas prone to thermokarst formation in the future.

Author contributions. XP and GY designed the research and obtained funding. GY analyzed the data and prepared the data files. GY, WT, XL and XP conducted the field work. GY, XP, OWF, JL, CM, FN wrote the paper with input from the coauthors and coordinated the analysis and contributions from all coauthors. XP and GY contributed to the work equally and should be regarded as co-first authors.

Competing interests. The contact author has declared that neither they nor their coauthors have any competing interests.

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