

Dataset of Stable Isotopes of Precipitation in the Eurasian Continent

Longhu Chen^{1,2,3, a}, Qinqin Wang^{1,2,3, a}, Guofeng Zhu^{1,2,3*}, Xinrui Lin^{1,2,3}, Dongdong Qiu^{1,2,3}, Yinying Jiao^{1,2,3}, Siyu Lu^{1,2,3}, Rui Li^{1,2,3}, Gaojia Meng^{1,2,3}, Yuhao Wang^{1,2,3}

¹ School of Geography and Environment Science, Northwest Normal University, Lanzhou 730070, Gansu, China

² Shiyang River Ecological Environment Observation Station, Northwest Normal University, Lanzhou 730070, Gansu, China

³ Key Laboratory of Resource Environment and Sustainable Development of Oasis, Lanzhou 730000, China

^a These authors contributed equally to this work and should be considered co-first authors.

Correspondence to: zhugf@nwnu.edu.cn

Abstract: Stable isotopes in precipitation can effectively reveal the process of atmospheric water circulation, serving as an effective tool for hydrological and water resources research, climate change, and ecosystem studies. The scarcity of stable isotope data in precipitation has hindered comprehension of regional hydrology, climate, and ecology due to discontinuities on a temporal scale and unevenness on a spatial scale. To this end, we collated stable hydrogen and oxygen isotope data in precipitation from 842 stations in Eurasia from 1961 to 2022, totalling 51,752 data records. Stable isotopes in precipitation across various regions of Eurasia, as a whole, decrease with increasing latitude and distance from the coast. In the summer, stable isotopes in precipitation are relatively enriched, while in winter, they are relatively depleted. In recent decades, the stable isotope values of Eurasian precipitation show an overall trend of increasing variation with the advancement of years, which is associated with global warming. Geographical location, underlying surface conditions, seasons, and atmospheric circulation are all factors that determine the characteristics of stable isotopes in precipitation. The dataset of stable isotopes in Eurasian precipitation provides a powerful tool for understanding changes in regional atmospheric water circulation and assists in conducting hydrological, meteorological, and ecological studies in related regions.

Keywords: Eurasian Continent, Climate Change, Stable Isotopes in Precipitation, Atmospheric Circulation, Dataset.

29 **1. Introduction**

30 In recent years, the impacts of global climate change have become increasingly severe,
31 particularly the significant increase in the frequency of various types of extreme weather and
32 climate events (Faranda et al., 2023; Liu et al., 2022; Zhang et al., 2016). The World
33 Meteorological Organization's 2022 report on the state of the climate in Asia shows that the
34 rate of warming in Asia is higher than the global average, with droughts, floods, and
35 heatwaves affecting most parts of the world (State of the Climate in Asia 2022). Severe
36 fluctuations in climatic elements can alter water circulation processes, affect regional climate
37 change, and even change the evolutionary patterns of ecological environments. Among these,
38 stable isotopes in precipitation are an excellent comprehensive tracer, playing an important
39 role in revealing water cycle processes, climate change information, and mechanisms of water
40 resource use in ecosystems (Bowen et al., 2019; Wang et al., 2022). Therefore, in the face of
41 increasingly complex climate conditions, we need more comprehensive data on stable
42 isotopes in precipitation at various spacetime scales to help understand climate change
43 phenomena.

44 Stable isotopes in precipitation serve as a crucial medium connecting the hydrological
45 and climatic systems. Precipitation, being both a product of the climate system and a primary
46 source for the hydrological system (Sun et al., 2018), plays a pivotal role. Additionally, stable
47 isotope fractionation accompanying the water cycle not only carries rich climate information
48 throughout its variations but also facilitates the tracing of contributions to various surface
49 water bodies (Hao et al., 2019; Ren et al., 2017; Shi et al., 2022). Although stable isotopes in
50 precipitation ($\delta^2\text{H}$ and $\delta^{18}\text{O}$) constitute a small proportion in natural water bodies, they exhibit
51 sensitivity to changes in climatic factors (Craig, 1961; Dansgaard, 1964). The quantification
52 of precipitation stable isotopes, influenced by factors such as temperature, precipitation, wind
53 speed, relative humidity, and water vapour sources (Gat, 1996; Jiao et al., 2020), deepens our
54 procedural understanding of the water cycle. This quantification provides relevant information
55 about water vapour transport processes and precipitation formation (Kathayat et al., 2021),
56 determination of the proportions of different types of precipitation (Aggarwal et al., 2016),
57 and comprehension of the mechanisms behind extreme events (Sun et al., 2022), offering
58 robust evidence to explore the inherent mechanisms of meteorological events and climate

59 change processes. Water recovery is a significant component of land water flux (Jasechko et
60 al., 2013), but its direct measurement still faces numerous challenges. Deuterium excess
61 (d-excess): $\delta^2\text{H} = 8 \times \delta^{18}\text{O}$, a stable isotope quantity sensitive to water recovery effects,
62 remains constant throughout the entire process from water vapor evaporation into the
63 atmosphere to final condensation and rain formation (Merlivat and Jouzel, 1979). Therefore,
64 in current water recovery quantification efforts, precipitation stable isotopes are a primary
65 means (Cropper et al., 2021; Zhang et al., 2021a). $\delta^2\text{H}$ and $\delta^{18}\text{O}$, as important climate tracers,
66 are also employed in reconstructing continental paleoclimate. The accurate understanding of
67 precipitation stable isotopes' response to modern climate lays the foundation for paleoclimate
68 reconstruction. On the other hand, using general atmospheric circulation models to simulate
69 isotope circulation is a major method for comparing isotope distributions in precipitation
70 under both modern and ancient conditions (Joussaume et al., 1984; Brady et al., 2019).
71 Simultaneously, the comparison between simulated and observed precipitation stable isotopes
72 provides valuable validation for the physical components of atmospheric circulation models
73 (Joussaume et al., 1984; Ruan et al., 2019). In conclusion, the comprehensive data on stable
74 isotopes in precipitation offer more detailed information about the climate and hydrological
75 systems.

76 In 1961, the International Atomic Energy Agency (IAEA) and the World Meteorological
77 Organization (WMO) began establishing the Global Network for Isotopes in Precipitation
78 (GNIP), which is the world's primary observation system. To date, research on stable isotopes
79 in precipitation primarily relies on the GNIP database. However, GNIP's observations are very
80 unevenly distributed in time and space. Global and regional-scale research on stable isotopes
81 in precipitation mainly depends on model simulations. The relationship between predicted
82 data from models and actual measured data is "comparative" (Joussaume et al., 1984).
83 Although model simulations can compensate for the absence of measured data and are
84 particularly advantageous in revealing the operating mechanisms of large-scale climate
85 systems and water cycles, existing models for stable isotopes in precipitation are often
86 insufficiently accurate. They cannot check long-term trends or characteristics of interannual
87 variation. By integrating independent data to provide a higher density of data, it's possible to
88 enhance the precision of model simulations.

89 We have compiled stable isotopes in precipitation data from the Eurasian continent since
90 1961 with the aim of providing more comprehensive data support for the following research
91 areas:

92 Climate research: stable isotopes in precipitation exhibit geographical and seasonal
93 variations, which can be used to study climate change and the impact of solar radiation. By
94 comparing and analyzing the stable isotopes of precipitation in different regions of the
95 Eurasian continent, long-term climate trends can be revealed, such as changes in precipitation
96 distribution and the evolution of monsoon systems.

97 Earth system research: stable isotopes in precipitation are not only influenced by climate
98 and water cycle but also by geological and biological processes. By integrating precipitation
99 stable isotope data from the Eurasian continent, it is possible to investigate in-depth the
100 interactions between different components of the Earth system, such as the interaction
101 between the atmosphere and the ocean, and the water cycle in terrestrial ecosystems. This will
102 contribute to a better understanding of the functioning and changes of the Earth system.

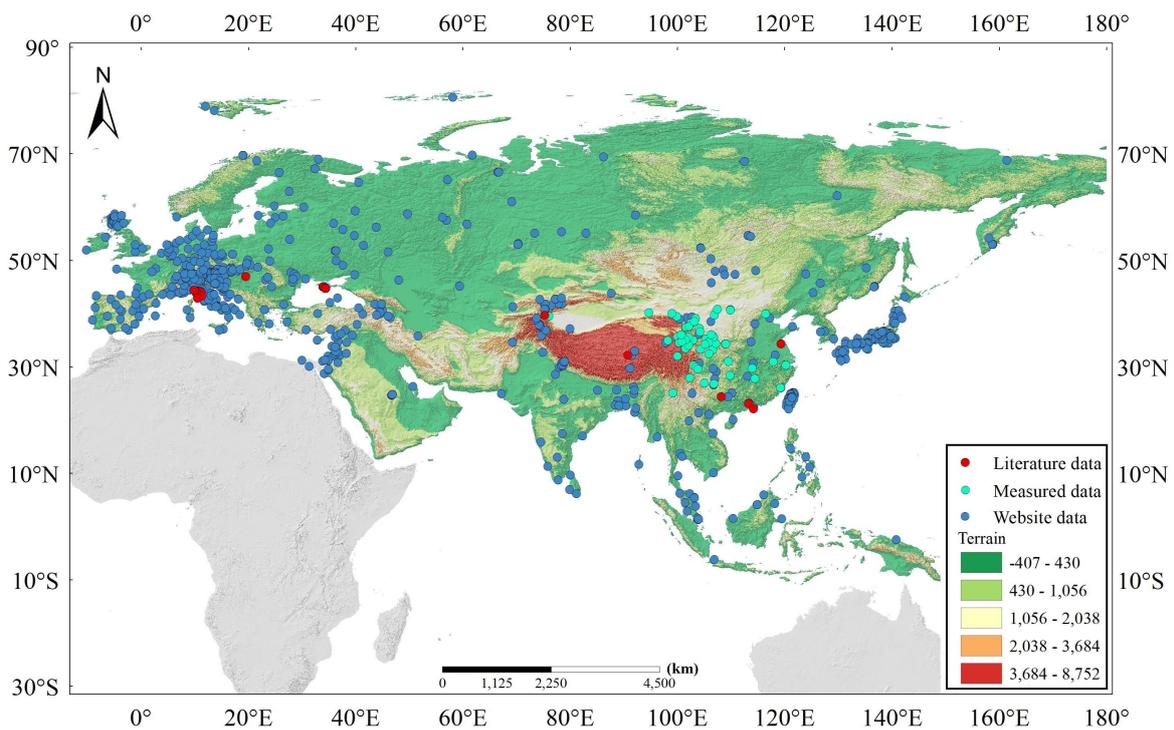
103 Water cycle research: stable isotopes in precipitation serve as important indicators of the
104 water cycle and can track the sources, evaporation, and precipitation processes of water. By
105 analyzing the spatial distribution and variations of precipitation stable isotopes on the
106 Eurasian continent, it is possible to understand the processes of water evaporation,
107 precipitation, and recycling, revealing the patterns of water resource distribution and changes.
108 This provides support for water resource management and hydrological modelling.

109 Paleoclimate Reconstruction: Well-established precipitation stable isotope observational
110 data are advantageous for validating paleoclimate models under modern conditions.
111 Simultaneously, they contribute to richer comparative data for stable isotopes in precipitation
112 collected in geological archives.

113 **2. Study area**

114 The Eurasian continent (10°45'N - 77°44'N, 9°30'W - 169°45'E) spans a vast territory,
115 with considerable variations in natural geographic conditions within the region (Fig.1).
116 Significant thermal differences between sea and land have given rise to a typical monsoon
117 climate system on the southeast coast, while interactions between Atlantic moisture and
118 planetary wind systems result in the west coast and wide inland areas being perennially

119 subject to westerly moisture. These two major systems play significant roles in global climate
 120 systems (Li et al., 2022; Wang et al., 2010). Moreover, the interactions across multiple heat
 121 zones with sea and land provide conditions conducive to a wide variety of climate types. The
 122 uplift of the Qinghai-Tibet plateau not only alters the climate patterns dominated by the
 123 planetary wind system on the Eurasian continent and the moisture movement paths in the
 124 Indian Ocean (An et al., 2001) but also changes the natural surface conditions, such as the
 125 numerous rivers including the Yangtze, Yellow, Ganges, and Mekong Rivers, which play a
 126 vital role in hydrological processes and human life. The plateau itself forms a relatively
 127 complete vertical ecological environment differentiation, enhancing the complexity of the
 128 natural environment on the Eurasian continent. Therefore, the research data and studies on
 129 climate environmental changes in Eurasia hold significant representativeness in addressing
 130 global changes.



131
 132 **Fig.1** Distribution map of precipitation stable isotope sampling sites in the Eurasian continent

133 **3. Data and methodology**

134 **3.1 Data sources and collection**

135 We have collected $\delta^{18}\text{O}$ and $\delta^2\text{H}$ stable isotope data from precipitation at 842 sa

136 mpling points across the Eurasian continent from 1961 to 2022(Supplement, Table S1).
137 The dataset includes both measured data and data collected from various sources. Th
138 e data collected are primarily from the Water Isotopes website (<https://wateriso.utah.edu>
139 /waterisotopes/index.html) and the Global Network of Stable Isotopes in Precipitation
140 (GNIP) operated by the International Atomic Energy Agency (IAEA). In this study, we
141 have compiled a total of 45,782 data records, including 3,676 records from literature
142 sources. The measured data were collected, analyzed, and organized at the Shiyang R
143 iver Basin Integrated Observation Station of Northwest Normal University in China, co
144 mprising 2,297 data records. Additionally, meteorological data used in this study are fr
145 om the CRU TS v. 4.07 dataset (Harris et al., 2020) and the NCEP-NCAR Reanalysis
146 1 dataset (<https://psl.noaa.gov/data/gridded/data.ncep.reanalysis.html>). As well as, As w
147 ell as, the global climate classification data of Köppen (Beck et al., 2018) (Suppleme
148 nt, S2).

149 3.2 Data processing steps and quality control

150 Data Collection: The data collected includes a variety of issues such as missing values,
151 outliers, and duplicates, as well as gaps in dates and missing or incorrect latitude and
152 longitude information. Therefore, the collected raw data underwent preprocessing and data
153 cleaning. Missing data was interpolated, entries that could not be completed were removed,
154 and duplicate data was eliminated.

155 Measured Data: Standard rain gauges were used to collect precipitation samples. After
156 each precipitation event, the collected samples were immediately transferred into 100ml
157 high-density sample bottles. To prevent data errors caused by evaporation, the collected water
158 samples were stored in a refrigerator at a temperature of approximately 4°C. Prior to analysis,
159 the precipitation samples were naturally thawed at room temperature. Impurities were filtered
160 out using a 0.45µm filter membrane, and the samples were transferred to 2ml sample bottles.
161 Isotope values were measured using a liquid water isotope analyzer (DLT-100, Los Gatos
162 Research, USA). For any abnormal or values that did not pass the LWIA post-analysis
163 software check, parallel samples were selected for re-measurement to ensure data accuracy
164 (Zhu et al., 2022). The isotopic abundances of ¹⁸O and ²H were expressed using the δ notation

165 relative to the International Atomic Energy Agency (IAEA) Vienna Standard Mean Ocean
166 Water (V-SMOW) reference, following the equation:

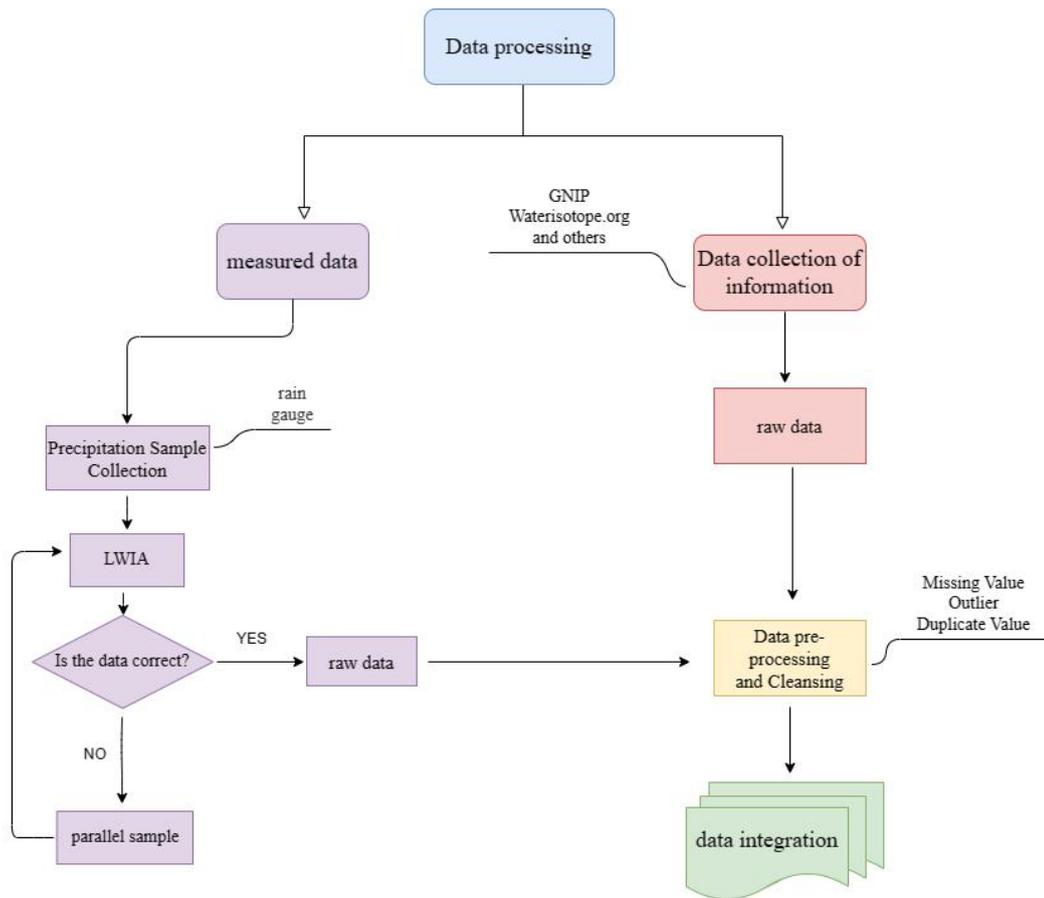
$$167 \quad \delta_{\text{sample}} (\text{‰}) = \left[\frac{R_{\text{sample}}}{R_{V\text{-}smow}} - 1 \right] \times 1000$$

168 Here, R represents the ratio of the heavier isotope to the lighter isotope (i.e., $^{18}\text{O}/^{16}\text{O}$ or
169 $^2\text{H}/^1\text{H}$). We used the International Atomic Energy Agency (IAEA) standard (V-SMOW2) to
170 validate our isotope measurements, ensuring comparability between isotopic measurements
171 across laboratories and instruments.

172 In 1982, Ferronsky VI and Polyakov VA conducted a study that found a general
173 distribution of $\delta^{18}\text{O}$ and $\delta^2\text{H}$ values in natural substances, indicating that the range of stable
174 isotope values for hydrogen and oxygen in atmospheric precipitation is typically -400‰ to
175 -30‰ and -60‰ to 10‰, respectively (Ferronsky VI et al., 1982). After data processing, the
176 data generally falls within the reasonable range.

177 In addition, we selected the two climatic zones with the most significant differences,
178 namely the tropical and polar zones. The reason for this choice is that the boundaries between
179 temperate, frigid, and arid zones are relatively unclear, with subtle changes in trends.
180 Mann-Kendall (MK) tests were conducted on the temporal variations of stable isotopes in
181 precipitation for both climatic zones (Fig.3). For the tropical climate (A), the stable isotopes
182 of precipitation ($\delta^2\text{H}$ and $\delta^{18}\text{O}$) exhibit multiple non-significant periods of abrupt changes.
183 There is a significant increasing trend from 1971 to 2005, followed by a non-significant
184 decreasing trend since 2009. Overall, the deuterium excess (d-excess) shows a non-significant
185 decreasing trend, but this trend has weakened after 1990. In the polar climate (E), there is a
186 significant increasing trend before 1973, followed by non-significant periods of both increase
187 and decrease after 1975. However, after 2010, a gradually significant increasing trend is
188 observed. Since 1985, the deuterium excess has undergone a non-significant decreasing
189 process, and after 2010, it gradually reaches a significant increasing trend. The uncertainty of
190 the tests is mainly attributed to the spatiotemporal distribution and volume of the data.

191



192
193
194

Fig.2 Flowchart of precipitation stable isotope dataset construction

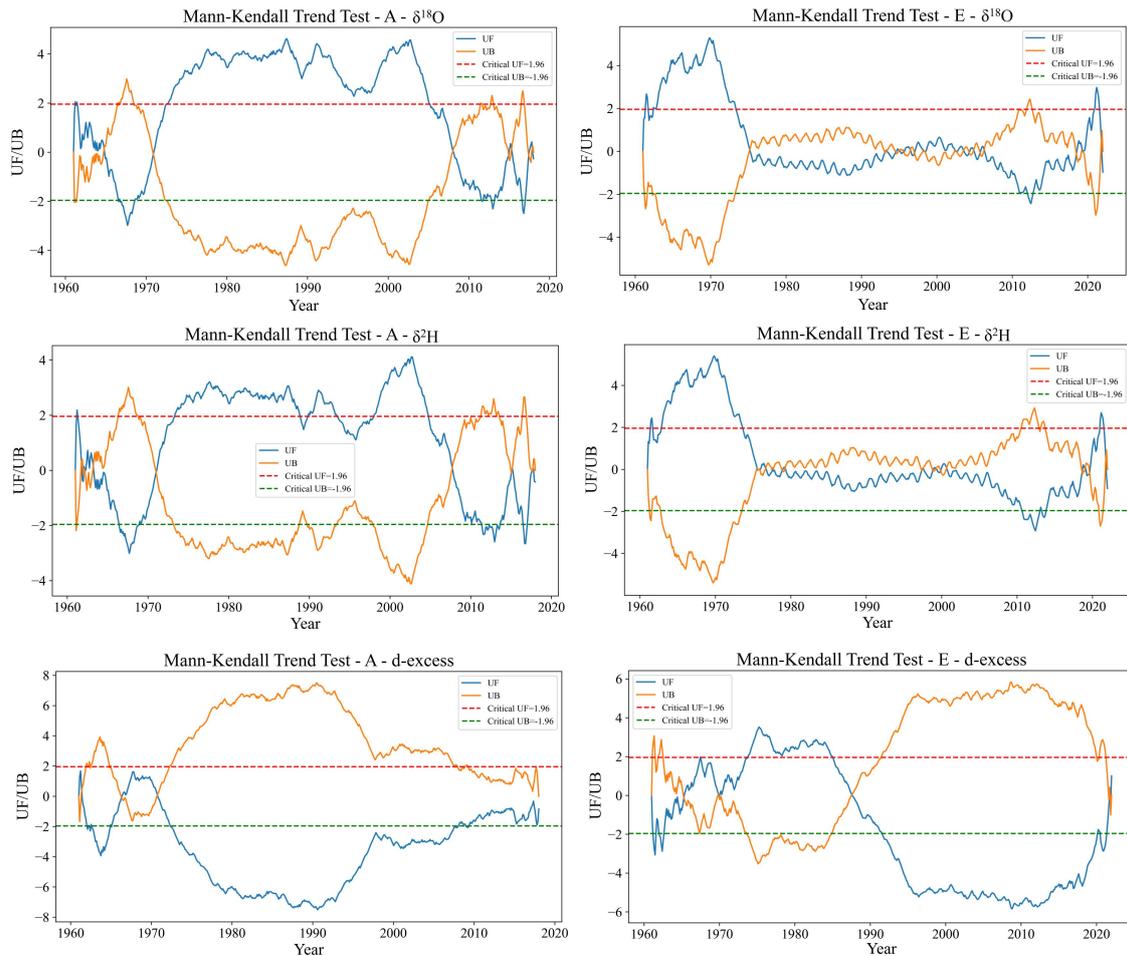


Fig.3 Time series MK test for temperate (C) and cold (D) climates

195
196
197

4. Results and discussion

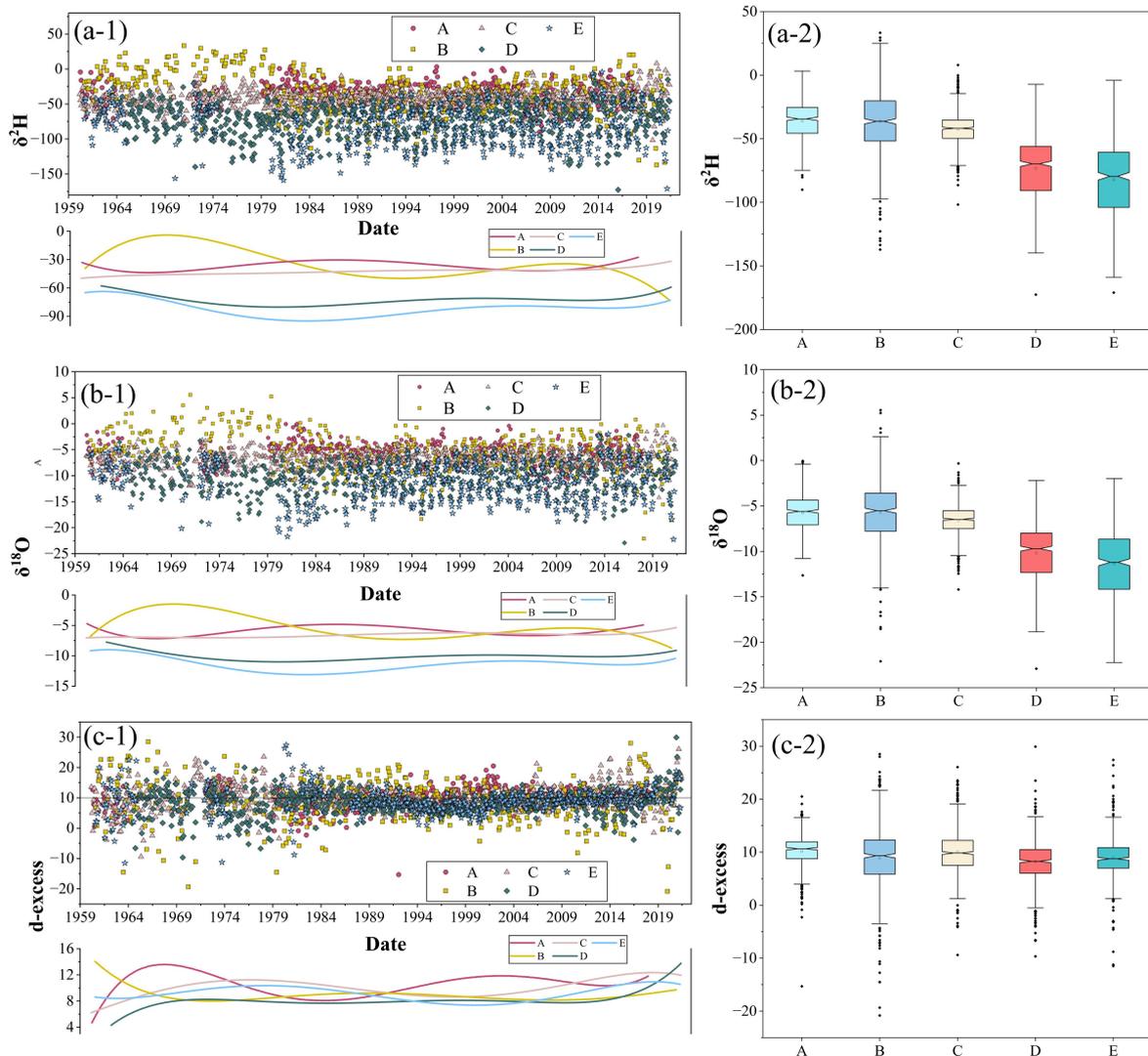
4.1 Temporal and Spatial Variation Characteristics of Precipitation Stable Isotopes

198
199
200 On a temporal scale, stable isotopes in precipitation exhibit pronounced seasonal
201 variations, with higher values during the summer and lower values during the winter (Figure
202 4). This is attributed to seasonal variations in evaporation caused by temperature changes,
203 resulting in the evaporative fractionation of stable isotopes in precipitation. Considering the
204 completeness of the time series and regional differences within the Eurasian continent, we
205 constructed a time series of precipitation stable isotopes based on the Köppen climate
206 classification "climate zones" The temporal changes in precipitation stable isotopes under
207 different climate types show significant differences. In tropical climates (A), the values of
208 precipitation stable isotopes are higher, with low values reflecting enhanced precipitation. The
209 "precipitation effect" in the Eurasian continent is particularly significant in tropical climates

210 (Tharammal et al., 2017), and the composition of precipitation stable isotopes reflects the
211 correlated changes between temperature and precipitation. However, the seasonal fluctuations
212 in tropical precipitation stable isotopes are minimal, and there is a fluctuating trend over
213 approximately 20 years. Most arid climates (B) and temperate climates (C) on the Eurasian
214 continent are under the influence of the westerly system. Before 1980, temperate climates
215 experienced significant fluctuations in precipitation stable isotopes, followed by a stable
216 period of about 30 years. After 2010, an unstable trend has become more pronounced,
217 reflecting an increase in extreme weather events (Yao et al., 2021; Zhang et al., 2012). In arid
218 climate regions, precipitation stable isotopes have undergone significant decreases. The
219 Central Asian arid region is a typical temperate arid region, and numerous studies have
220 pointed out a "warm and humid" trend in the climate of this region (Wang et al., 2020; Yan et
221 al., 2019). The strengthening of the West Pacific subtropical high, North American subtropical
222 high, and the Asian subtropical westerly jet is believed to increase precipitation in this region
223 (Chen et al., 2011). The enhancement of high-latitude water vapour transport is a major factor
224 influencing the increase in precipitation in the Central Asian arid region, which is also the
225 reason for the decreasing trend in deuterium excess (Fig. 4, c-1). Cold climates (D) and polar
226 climates (E) have the smallest values of precipitation stable isotopes, but they exhibit
227 significant differences on an annual scale and a gradually increasing trend on an interannual
228 scale. With global warming, high-latitude regions will provide more sources of water vapour
229 for the water cycle (Ding et al., 2017).

230 On a spatial scale, the topographic differences and latitude variations in the region are
231 the primary causes of spatial differences in stable isotopes in precipitation across the Eurasian
232 continent. The multi-year average values of $\delta^2\text{H}$ and $\delta^{18}\text{O}$ at different latitudes are as follows:
233 from 0° to 30°N , they are -30.20‰ and -5.99‰ , from 30° to 60°N , they are -58.94‰ and
234 -8.77‰ , and from 60° to 90°N , they are -92.98‰ and -12.69‰ . The Alps and the Tibetan
235 Plateau form regions of low precipitation stable isotopes that differ from those at the same
236 latitudes. The gradual uplift of the Tibetan Plateau's mountains leads to changes in the
237 atmospheric circulation patterns over a larger area, altering the inherent characteristics of
238 water vapour source regions, vapour transport paths, and precipitation stable isotope values.
239 The response of precipitation stable isotopes to the plateau's climate reflects changes in the

240 large-scale circulation state (Yao et al., 2013). The isotopic variations in the surrounding
 241 regions of the Alps reflect differences in water vapour sources due to regional topography
 242 (Natali et al., 2021; Rindsberger et al., 1983). Spatial variations in deuterium excess can
 243 effectively reflect differences in regional water vapour sources, with average values of
 244 approximately 10‰ for tropical and temperate climates. Cold climate regions have lower
 245 deuterium excess values, and due to the overlap of arid climates with other climate zones, the
 246 distribution range of deuterium excess values in arid climates is larger. Therefore, it can be
 247 hypothesized that if isotope-related variables (e.g., d-excess) are included in climate zone
 248 classification criteria, more climate zones influenced by circulation patterns could be
 249 identified.



250
 251

Fig.4 The time series variations of $\delta^2\text{H}$, $\delta^{18}\text{O}$, and d-excess in the Eurasian continent.

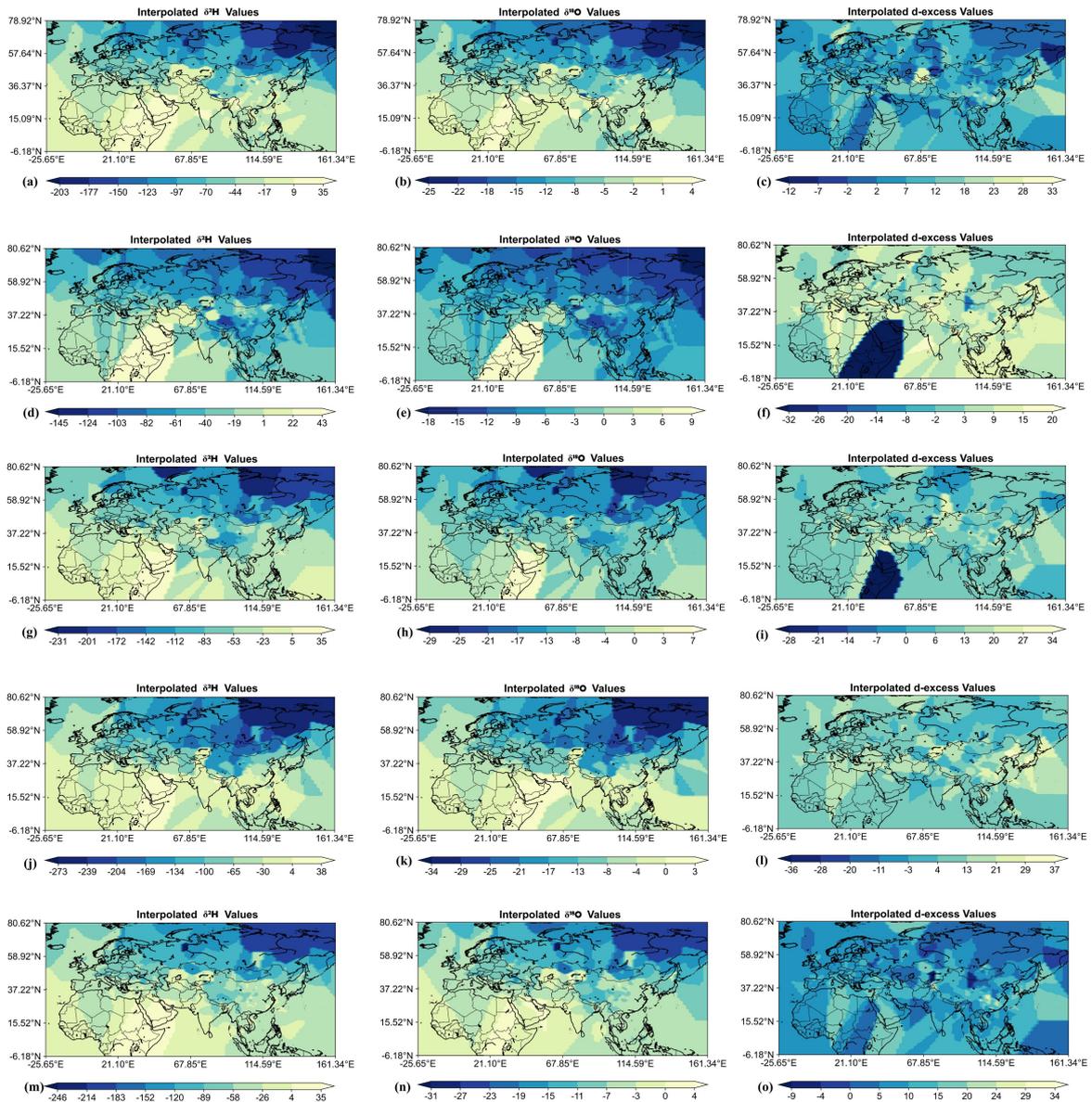


Fig.5 The spatial variations of $\delta^2\text{H}$, $\delta^{18}\text{O}$, and d-excess in the Eurasian continent. Panels (a), (b), and (c) display the spatial distribution of isotope values in the spring season. Panels (d), (e), and (f) show the spatial distribution of isotope values in the summer season. Panels (g), (h), and (i) present the spatial distribution of isotope values in the autumn season. Panels (j), (k), and (l) exhibit the spatial distribution of isotope values in the winter season. Panels (m), (n), and (o) display the spatial distribution of isotope values averaged over multiple years.

252

253 4.2 Seasonal changes in meteoric water line and precipitation stable isotopes

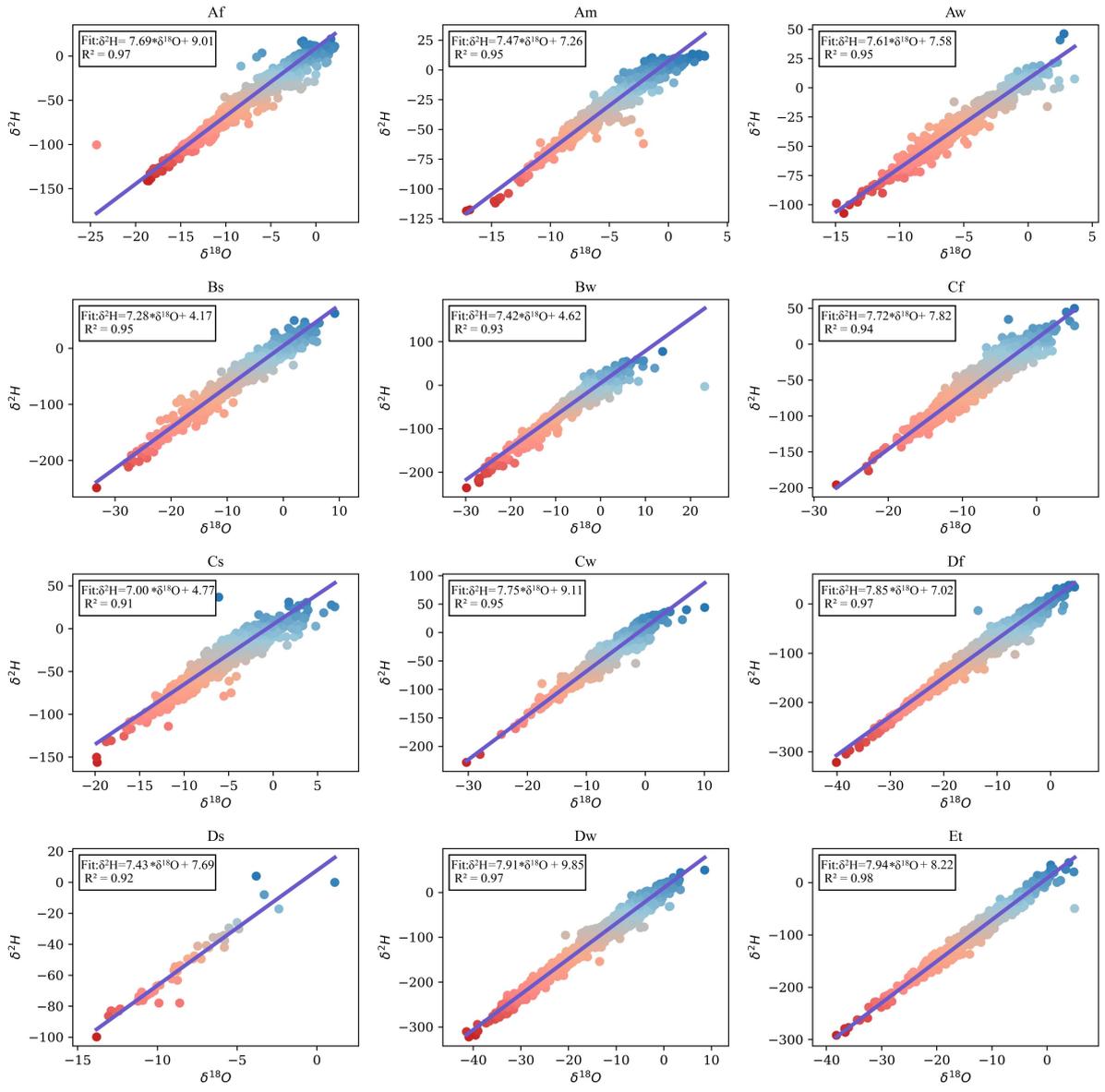
254 The temporal and spatial variations of stable isotopes in precipitation are greatly
 255 influenced by meteorological factors, and the changes in precipitation isotopes are consistent
 256 with the climatic regions. Therefore, based on the Köppen climate classification, we

257 performed climate zoning for stable isotopes in precipitation sites. We used the least squares
258 method to fit meteoric water line for different climate regions (Fig.6) and considered the
259 seasonal variations of precipitation stable isotopes in different climate regions (Fig.7). The
260 meteoric water line for different climate types indicate relatively small differences in various
261 climate precipitation amounts in tropical climates. The variations in the slope and intercept of
262 the meteoric water line are determined by the combined effects of precipitation and
263 temperature, with convective precipitation weakening the impact of the "temperature effect."
264 Intense convective rainfall and oceanic water vapour transport bring abundant precipitation to
265 tropical regions. The fractionation mechanisms and variations of precipitation stable isotopes
266 not only reveal the inherent patterns of weather pattern occurrence and development (Sun et
267 al., 2022) but also correlate weather patterns with supply sources, tracing the water sources of
268 surface water bodies (Scholl and Murphy, 2014; Anon, 2017). Stable isotopes in precipitation
269 in arid climates are influenced by secondary evaporation below clouds, and intense
270 unbalanced fractionation processes lead to relative enrichment of stable isotopes in
271 precipitation (Wang et al., 2021; Zhu et al., 2021). Water resources are the most limiting
272 factor for the ecological and social environment in arid climate regions (García-Ruiz et al.,
273 2011). Therefore, compared to other climate regions, water recovery becomes more critical.
274 Stable isotopes in precipitation can accurately quantify water recovery and effectively assess
275 the impact of evaporation on different water bodies in arid regions. The majority of the global
276 population is distributed in temperate regions. Therefore, with global temperature rise, the
277 climate change situation in temperate regions deserves more attention. In temperate climate
278 zones, the differences in stable isotope composition between different climate types become
279 more significant. In the Mediterranean region controlled by the Summer Dry Warm Climate,
280 the slope and intercept are the lowest, indicating that temperature rise dominates the
281 fractionation of hydrogen and oxygen stable isotopes in precipitation, and the region shows a
282 trend of aridification under long-term average conditions. The westerly system is the main
283 controlling circulation in this region, and the changes in precipitation stable isotopes reflect
284 the attenuation trend of mid-latitude westerly moisture inward migration (Zhu et al., 2023; Shi
285 et al., 2021). In polar climates, the atmospheric water line exhibits higher slope and intercept.
286 The influence of unbalanced fractionation processes after water vapour condensation in cloud

287 systems is relatively small, resulting in a slope close to 8.

288 The seasonal variation of hydrogen and oxygen stable isotopes in precipitation on the
289 Eurasian continent generally exhibits a pattern of higher values in summer and lower values
290 in winter (Fig.7)(Hydrogen isotopes ($\delta^2\text{H}$) in Supplement S3). However, there are still
291 significant differences under different climate zones. The seasonal differences in tropical
292 climates are less pronounced, with the Tropical Sparse Forest Climate (Aw) showing a
293 decrease and increase with the months, possibly due to an increase in precipitation. Temperate
294 and cold climates generally exhibit significant seasonal variations. The deuterium excess in
295 the Eurasian continent shows a lower pattern in summer and a higher pattern in winter,
296 indicating seasonal changes in water vapour sources and transport distances (Zhang et al.,
297 2021a). This overall suggests that the summer climate in Eurasia is more humid, while the
298 winter climate is drier. Deuterium excess usually indicates the degree of imbalance in
299 seawater sources during their evaporation process, and it typically depends only on the
300 environmental conditions of the evaporation source. Compared to $\delta^2\text{H}$ and $\delta^{18}\text{O}$, deuterium
301 excess displays a more stable pattern and is distributed around the global average (10‰). The
302 westerly and monsoon systems are the primary atmospheric circulation systems over the
303 Eurasian continent, carrying water vapour from the ocean inland and gradually weakening.
304 This indicates that the humidity in the vast region of Eurasia is strongly influenced by ocean
305 water vapour. Ocean conditions and large-scale atmospheric circulation changes can have
306 profound effects on the climate environment of the Eurasian continent.

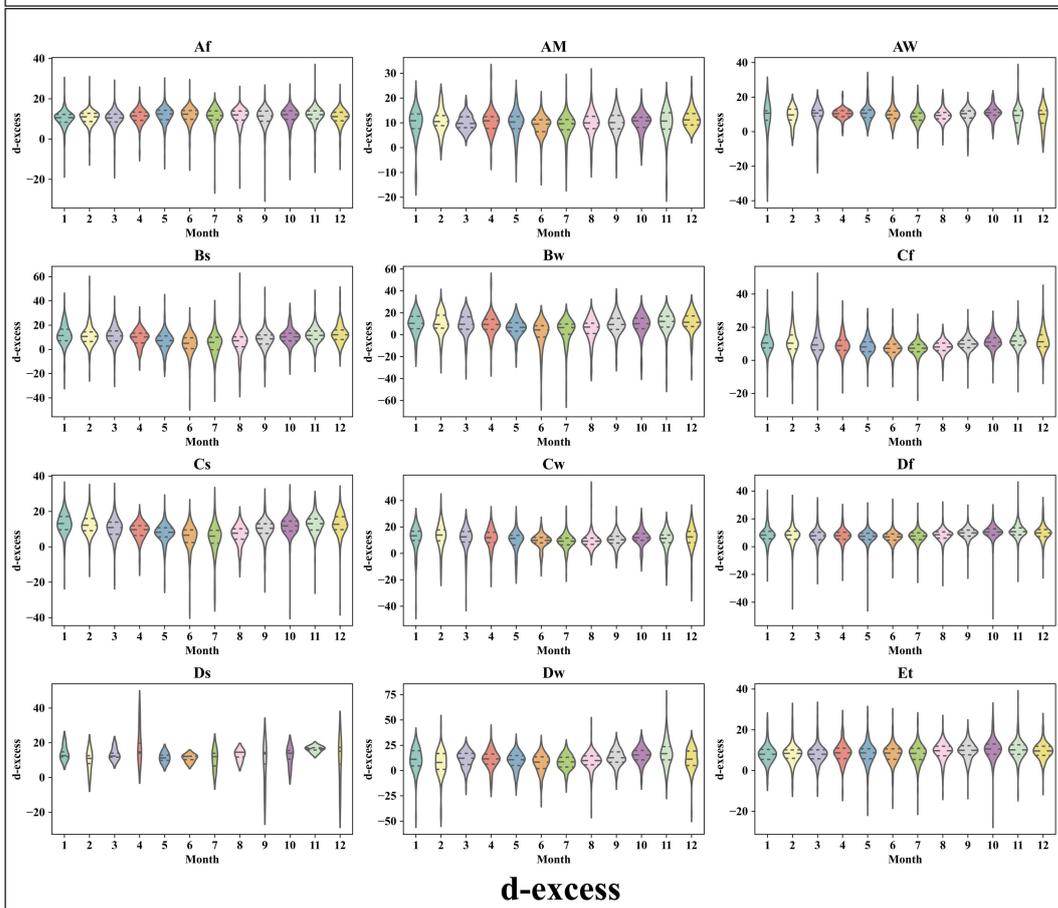
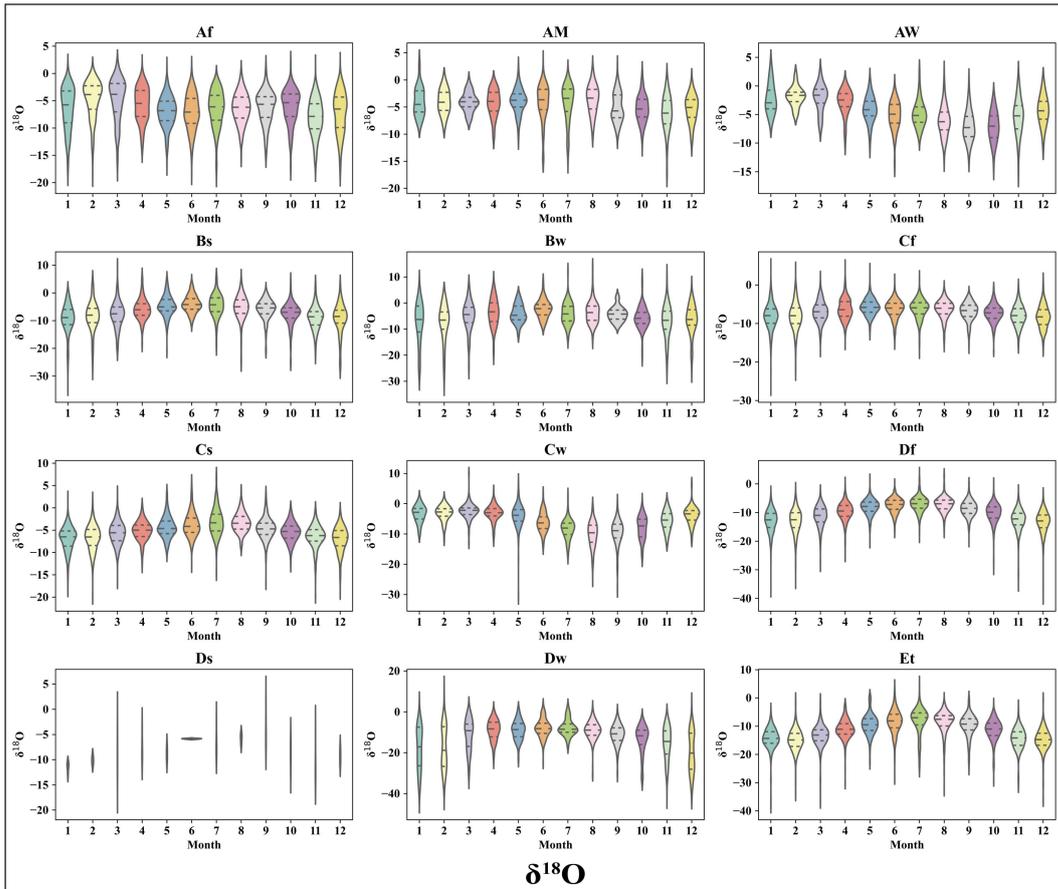
307 The differences in precipitation stable isotopes among different climate types are not
308 only responses to different climate characteristics but also provide effective tools for a deeper
309 understanding of the process, climate change mechanisms, water vapour transport between
310 land and sea, and supply relationships between water bodies. The precipitation stable isotopes
311 dataset we have constructed for the Eurasian continent can be combined with traditional
312 meteorological data to provide more information for climate and hydrological research.



313

314

Fig.6 Different meteoric water lines in various climate zones.



316 **Fig.7** Seasonal Distribution and Variations of Stable Isotopes in Precipitation ($\delta^{18}\text{O}$, d-excess)

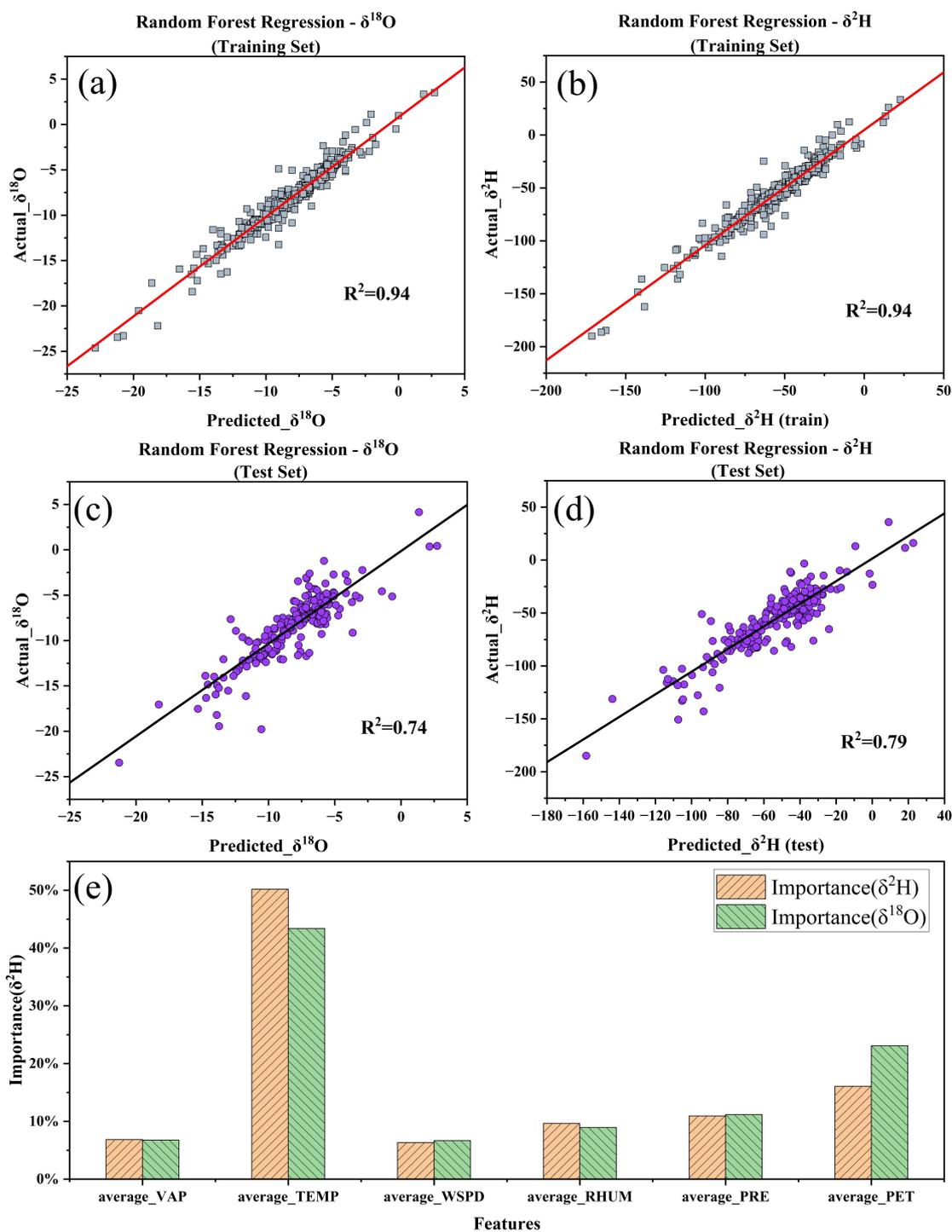
317 4.3 Drivers of stable isotope variation in precipitation in Eurasia

318 Meteorological variables accompany the fractionation process of stable hydrogen and
319 oxygen isotopes in precipitation, impacting the composition of stable isotopes (Sun et al.,
320 2019). We utilized a random forest regression model to assess the importance of
321 meteorological variables in the Eurasian continent on stable isotopes. Random forest
322 regression is a non-parametric method used to solve prediction problems. It predicts
323 regression problems based on the average results of random decision trees, which use
324 bootstrapping to eliminate the possibility of overfitting (Erdélyi et al., 2023). The random
325 forest regression analysis of the fitted stable isotopes of hydrogen and oxygen showed good
326 goodness of fit for both the training and testing sets, indicating that temperature, precipitation,
327 potential evapotranspiration, vapour pressure, wind speed, and relative humidity have a high
328 explanatory power for stable isotopes of hydrogen and oxygen (Fig.8). The results of
329 cross-validation for the model indicate superior predictive performance for the target variable
330 $\delta^{18}\text{O}$ compared to the target variable $\delta^2\text{H}$, as reflected in the smaller root mean square error
331 (RMSE) and mean absolute error (MAE) for $\delta^{18}\text{O}$ (Supplement, Table S3). The composition
332 of stable isotopes in precipitation is greatly influenced by meteorological variables. Among
333 the six variables considered, temperature has the strongest explanatory power for the variation
334 of stable isotopes of hydrogen and oxygen, and potential evapotranspiration also has a
335 relatively strong explanatory ability, indicating that temperature change primarily drives the
336 variation of stable isotopes in precipitation in the Eurasian continent. The relative humidity is
337 the ratio of actual vapour pressure to saturated vapour pressure, but there is a significant
338 difference in the explanatory power of vapour pressure and relative humidity on stable
339 isotopes. Vapour pressure has a wider range of variation in the atmosphere, thus it may have
340 greater variability in the regression model, leading to a larger impact on predicting stable
341 isotopes in precipitation. Relative humidity, on the other hand, is a relative indicator with a
342 relatively smaller range of variation, so it may have a weaker predictive ability for stable
343 isotopes in precipitation in the regression model. The driving factors for the variation of stable
344 isotopes in precipitation in the Eurasian continent include climate change, seasonal variations,
345 topography and landforms, as well as water cycle processes, which collectively influence the

346 isotopic composition of precipitation. Atmospheric circulation directly affects the source of
347 water vapor and the path of moisture, while other factors primarily influence the composition
348 of stable isotopes in precipitation by altering temperature. For example, potential
349 evapotranspiration plays a crucial role in explaining the variation of stable isotopes in
350 precipitation. However, the control of meteorological variables on stable isotopes in
351 precipitation varies between regions. Studies on two precipitation stations in Crimea have
352 shown weak correlations between temperature, precipitation, and stable isotopes in
353 precipitation. The complex natural environment determines that no single factor has a
354 dominant control over the stable isotopes in precipitation in that region, and the composition
355 of stable isotopes in precipitation is influenced by both local and distant factors (Dublyansky
356 et al., 2018). In the eastern coastal region of China, the relative enrichment of stable isotopes
357 in precipitation is due to the proximity to the evaporative source of the ocean, leading to an
358 increased abundance of heavy isotopes (Zhang et al., 2021b). In the arid region of central Asia,
359 there is a strong correlation between stable isotopes in precipitation and temperature, and the
360 enrichment or depletion of stable isotopes in precipitation reflects the trend of temperature
361 change (Zhu et al., 2023). In summary, the meteorological control factors of the composition
362 of stable isotopes in precipitation vary in different regions. There is a strong relationship
363 between stable isotopes in precipitation and meteorological variables, and stable hydrogen
364 and oxygen isotopes may be considered essential climate response variables, which will
365 contribute to describing the hydrological cycle and better predicting the response of future
366 climate change and ecosystem changes.

367 Stable isotopes in precipitation, serving as indicators of climate and environment, play a
368 unique role in enhancing the process-oriented understanding of extreme weather events and
369 exploring hydrological connections between different water bodies. However, a limitation
370 remains in the insufficient observation of stable isotopes in precipitation. Therefore, isotope
371 atmospheric circulation models based on physical mechanisms have been widely applied to
372 predict stable isotopes in water (Risi et al., 2012; Bowen et al., 2019). Physical models with
373 different driving mechanisms can meet various usage needs, including paleoclimate
374 reconstruction. For example, CAM3 simulation outputs precipitation oxygen isotope data (Lin
375 et al., 2024). Machine learning is a novel approach for predicting stable isotopes in

376 precipitation, and European simulation practices indicate that oxygen isotope simulations
377 have shown good results, while simulations for hydrogen isotopes remain challenging (Nelson
378 et al., 2021). In general, uncertainties in both physical models and machine learning need
379 continuous improvement and refinement through real-world data. Additionally, an accurate
380 understanding of the influencing factors of stable isotopes in precipitation is fundamental for
381 achieving successful predictions through machine learning.



382
 383 **Fig.8** Results of Random Forest Regression Analysis for $\delta^2\text{H}$ and $\delta^{18}\text{O}$ in Relation to Meteorological
 384 Variables.a: Regression results for the training set of $\delta^{18}\text{O}$, b: Regression results for the training set of $\delta^2\text{H}$,
 385 c: Regression results for the testing set of $\delta^{18}\text{O}$, d: Regression results for the testing set of $\delta^2\text{H}$, e:
 386 Importance of meteorological variables for $\delta^2\text{H}$ and $\delta^{18}\text{O}$.

387 **5. Summary and outlook**

388 Stable isotopes in precipitation play a crucial role in both the climate and hydrological

389 systems, exhibiting sensitivity to variations in both time and space. Research indicates
390 significant differences in isotopic values between summer and winter, correlating with
391 seasonal changes in temperature and evaporation. The temporal and spatial variations of
392 precipitation stable isotopes vary significantly across different climate types, reflecting the
393 influence of climate characteristics on isotopic distribution. Terrain and latitude differences
394 are the primary reasons for spatial variations in stable isotopes in precipitation.
395 Meteorological factors have a notable impact on precipitation stable isotopes, as evidenced by
396 the meteoric water line in different climate types, revealing the influence of climate on
397 isotopic fractionation. Observations of precipitation stable isotopes contribute to
398 understanding weather patterns, water vapour sources, and transport pathways, providing
399 important insights into stable isotope variations in arid climates. The integrated dataset of
400 stable isotopes in precipitation from the Eurasian continent that we have compiled can offer
401 more detailed climate and hydrological information. However, future research efforts should
402 focus on improving observational data for Stable isotopes in precipitation. The uncertainties
403 in physical models and machine learning methods need refinement through additional
404 real-world data to enhance the accuracy of predicting precipitation stable isotopes.

405

406 **Data Availability**

407 Zhu, Guofeng (2024), “Dataset of Stable Isotopes of Precipitation in the Eurasian
408 Continent”, Mendeley Data, V2, doi: 10.17632/rbn35yrbd2.2

409 **Author Contribution Statement**

410 Longhu Chen: Conceptualization and Writing-Original draft preparation; Qinqin Wang:
411 Writing and Data processing; Guofeng Zhu: Writing review and editing; Xinrui Lin:
412 Modification; Dongdong Qiu: Modification; Yinying Jiao: data processing; Siyu Lu:
413 Experiment; Rui Li: Methodology; Gaojia Meng: Visualization; Yuhao Wang: Visualization.

414 **Declaration of Interest Statement**

415 The authors declare that they have no known competing financial interests or personal
416 relationships that could have appeared to influence the work reported in this paper.

417 **Acknowledgements**

418 This research was financially supported by the National Natural Science Foundatio

419 n of China(42371040, 41971036), Key Natural Science Foundation of Gansu Province
420 (23JRRA698), Key Research and Development Program of Gansu Province(22YF7NA1
421 22), Cultivation Program of Major key projects of Northwest Normal University(NWN
422 U-LKZD-202302), Oasis Scientific Research achievements Breakthrough Action Plan Pr
423 oject of Northwest Normal University(NWNU-LZKX-202303). The authors thank their
424 Northwest Normal University colleagues for their help in fieldwork, laboratory analysis,
425 and data processing.

426

427 **Reference**

428 Aggarwal, P. K., Romatschke, U., Araguas-Araguas, L., Belachew, D., Longstaffe, F. J.,
429 Berg, P., Schumacher, C., and Funk, A.: Proportions of convective and stratiform
430 precipitation revealed in water isotope ratios, *Nature Geosci*, 9, 624–629, [https://](https://doi.org/10.1038/ngeo2739)
431 doi.org/10.1038/ngeo2739, 2016.

432 Anon: Precipitation stable isotope records from the northern Hengduan Mountains in C
433 hina capture signals of the winter India–Burma Trough and the Indian Summer M
434 onsoon, *Earth and Planetary Science Letters*, 477, 123–133, [https://doi.org/10.1016/](https://doi.org/10.1016/j.epsl.2017.08.018)
435 [j.epsl.2017.08.018](https://doi.org/10.1016/j.epsl.2017.08.018), 2017.

436 Beck, H. E., Zimmermann, N. E., McVicar, T. R., Vergopolan, N., Berg, A., and Woo
437 d, E. F.: Present and future Köppen-Geiger climate classification maps at 1-km re
438 solution, *Sci. Data*, 5, 180214, <https://doi.org/10.1038/sdata.2018.214>, 2018.

439 Bowen, G. J., Cai, Z., Fiorella, R. P., and Putman, A. L.: Isotopes in the Water Cycle:
440 Regional- to Global-Scale Patterns and Applications, *Annu. Rev. Earth Planet. Sc*
441 *i.*, 47, 453–479, <https://doi.org/10.1146/annurev-earth-053018-060220>, 2019.

442 Brady, E., Stevenson, S., Bailey, D., Liu, Z., Noone, D., Nusbaumer, J., Otto-Bliesner,
443 B. L., Tabor, C., Tomas, R., Wong, T., Zhang, J., and Zhu, J.: The Connected I
444 sotopic Water Cycle in the Community Earth System Model Version 1, *Journal of*
445 *Advances in Modeling Earth Systems*, 11, 2547–2566, <https://doi.org/10.1029/2019>
446 [MS001663](https://doi.org/10.1029/2019MS001663), 2019.

447 Chen, F., Huang, W., Jin, L., Chen, J., and Wang, J.: Spatiotemporal precipitation vari
448 ations in the arid Central Asia in the context of global warming, *Sci. China Eart*

449 h Sci., 54, 1812–1821, <https://doi.org/10.1007/s11430-011-4333-8>, 2011.

450 Craig, H.: Isotopic Variations in Meteoric Waters, *Sci. New Ser.*, 133, 1702–1703, 196

451 1.

452 Cropper, S., Solander, K., Newman, B. D., Tuinenburg, O. A., Staal, A., Theeuwens, J.

453 J. E., and Xu, C.: Comparing deuterium excess to large-scale precipitation recycl

454 ing models in the tropics, *npj Clim Atmos Sci*, 4, 1–9, <https://doi.org/10.1038/s41>

455 612-021-00217-3, 2021.

456 Czuppon, G., Bottyán, E., Kristóf, E., Weidinger, T., Haszpra, L., and Kármán, K.: Sta

457 ble isotope data of daily precipitation during the period of 2013–2017 from K-pus

458 zta (regional background monitoring station), Hungary, *Data Brief*, 36, 106962, <https://doi.org/10.1016/j.dib.2021.106962>, 2021.

459

460 Dansgaard, W.: Stable isotopes in precipitation, *Tellus*, 16, 436–468, <https://doi.org/10.3>

461 402/tellusa.v16i4.8993, 1964.

462 Ding, Q., Schweiger, A., L’Heureux, M., Battisti, D. S., Po-Chedley, S., Johnson, N.

463 C., Blanchard-Wrigglesworth, E., Harnos, K., Zhang, Q., Eastman, R., and Steig,

464 E. J.: Influence of high-latitude atmospheric circulation changes on summertime A

465 rctic sea ice, *Nat. Clim. Change*, 7, 289–295, <https://doi.org/10.1038/nclimate3241>,

466 2017.

467 Dublyansky, Y. V., Klimchouk, A. B., Tokarev, S. V., Amelichev, G. N., Langhamer,

468 L., and Spötl, C.: Stable isotopic composition of atmospheric precipitation on the

469 Crimean Peninsula and its controlling factors, *J. Hydrol.*, 565, 61–73, <https://doi.or>

470 [g/10.1016/j.jhydrol.2018.08.006](https://doi.org/10.1016/j.jhydrol.2018.08.006), 2018.

471 Erdélyi, D., Hatvani, I. G., Jeon, H., Jones, M., Tyler, J., and Kern, Z.: Predicting spa

472 tial distribution of stable isotopes in precipitation by classical geostatistical- and

473 machine learning methods, *J. Hydrol.*, 617, 129129, <https://doi.org/10.1016/j.jhydrol.>

474 [2023.129129](https://doi.org/10.1016/j.jhydrol.2023.129129), 2023.

475 Faranda, D., Messori, G., Jezequel, A., and Vrac, M.: Atmospheric circulation compou

476 nds anthropogenic warming and impacts of climate extremes in Europe, *Proc. Nat*

477 *l. Acad. Sci.*, 120, e2214525120, <https://doi.org/10.1073/pnas.2214525120>, 2023.

478 García-Ruiz, J. M., López-Moreno, J. I., Vicente-Serrano, S. M., Lasanta-Martínez, T.,

479 and Beguería, S.: Mediterranean water resources in a global change scenario, *Eart*
480 *h-Science Reviews*, 105, 121–139, <https://doi.org/10.1016/j.earscrev.2011.01.006>, 20
481 11.

482 Gat, J. R.: OXYGEN AND HYDROGEN ISOTOPES IN THE HYDROLOGIC CYCL
483 E, 1996.

484 Hao, S., Li, F., Li, Y., Gu, C., Zhang, Q., Qiao, Y., Jiao, L., and Zhu, N.: Stable isot
485 ope evidence for identifying the recharge mechanisms of precipitation, surface wat
486 er, and groundwater in the Ebinur Lake basin, *Science of The Total Environment*,
487 657, 1041–1050, <https://doi.org/10.1016/j.scitotenv.2018.12.102>, 2019.

488 Harris, I., Osborn, T. J., Jones, P., and Lister, D.: Version 4 of the CRU TS monthly
489 high-resolution gridded multivariate climate dataset, *Sci. Data*, 7, 109, [https://doi.or](https://doi.org/10.1038/s41597-020-0453-3)
490 [g/10.1038/s41597-020-0453-3](https://doi.org/10.1038/s41597-020-0453-3), 2020.

491 Jasechko, S., Sharp, Z. D., Gibson, J. J., Birks, S. J., Yi, Y., and Fawcett, P. J.: Terre
492 strial water fluxes dominated by transpiration, *Nature*, 496, 347–350, [https://doi.org](https://doi.org/10.1038/nature11983)
493 [/10.1038/nature11983](https://doi.org/10.1038/nature11983), 2013.

494 Jiao, Y., Liu, C., Liu, Z., Ding, Y., and Xu, Q.: Impacts of moisture sources on the t
495 emporal and spatial heterogeneity of monsoon precipitation isotopic altitude effects,
496 *J. Hydrol.*, 583, 124576, <https://doi.org/10.1016/j.jhydrol.2020.124576>, 2020.

497 Joussaume, S., Sadourny, R., and Jouzel, J.: A general circulation model of water isot
498 ope cycles in the atmosphere, *Nature*, 311, 24–29, <https://doi.org/10.1038/311024a0>,
499 1984.

500 Kathayat, G., Sinha, A., Tanoue, M., Yoshimura, K., Li, H., Zhang, H., and Cheng, H.:
501 Interannual oxygen isotope variability in Indian summer monsoon precipitation re
502 flects changes in moisture sources, *Commun Earth Environ*, 2, 96, [https://doi.org/1](https://doi.org/10.1038/s43247-021-00165-z)
503 [0.1038/s43247-021-00165-z](https://doi.org/10.1038/s43247-021-00165-z), 2021.

504 Li, G., Wang, X., Zhang, X., Yan, Z., Liu, Y., Yang, H., Wang, Y., Jonell, T. N., Qia
505 n, J., Gou, S., Yu, L., Wang, Z., and Chen, J.: Westerlies-Monsoon interaction dri
506 ves out-of-phase precipitation and asynchronous lake level changes between Centra
507 l and East Asia over the last millennium, *CATENA*, 218, 106568, [https://doi.org/1](https://doi.org/10.1016/j.catena.2022.106568)
508 [0.1016/j.catena.2022.106568](https://doi.org/10.1016/j.catena.2022.106568), 2022.

509 Lin, F., Zhang, Q., Sinha, A., Wang, Z., Axelsson, J., Chen, L., Wang, T., and Tan,
510 L.: Seasonal to decadal variations of precipitation oxygen isotopes in northern Chi
511 na linked to the moisture source, *npj Clim Atmos Sci*, 7, 1–9, [https://doi.org/10.1](https://doi.org/10.1038/s41612-024-00564-x)
512 038/s41612-024-00564-x, 2024.

513 Liu, Y., Cai, W., Lin, X., and Li, Z.: Increased extreme swings of Atlantic intertropica
514 l convergence zone in a warming climate, *Nat. Clim. Change*, 12, 828–833, <https://doi.org/10.1038/s41558-022-01445-y>, 2022.

516 Merlivat, L. and Jouzel, J.: Global climatic interpretation of the deuterium-oxygen 18 r
517 elationship for precipitation, *Journal of Geophysical Research: Oceans*, 84, 5029–5
518 033, <https://doi.org/10.1029/JC084iC08p05029>, 1979.

519 Natali, S., Baneschi, I., Doveri, M., Gianecchini, R., Selmo, E., and Zanchetta, G.:
520 Meteorological and geographical control on stable isotopic signature of precipitatio
521 n in a western Mediterranean area (Tuscany, Italy): Disentangling a complex signa
522 l, *J. Hydrol.*, 603, 126944, <https://doi.org/10.1016/j.jhydrol.2021.126944>, 2021.

523 Nelson, D. B., Basler, D., and Kahmen, A.: Precipitation isotope time series prediction
524 s from machine learning applied in Europe, *Proc. Natl. Acad. Sci.*, 118, e2024107
525 118, <https://doi.org/10.1073/pnas.2024107118>, 2021.

526 State of the Climate in Asia 2022: [https://library.wmo.int/records/item/66314-state-of-the-](https://library.wmo.int/records/item/66314-state-of-the-climate-in-asia-2022)
527 [climate-in-asia-2022](https://library.wmo.int/records/item/66314-state-of-the-climate-in-asia-2022), last access: 21 January 2024.

528 Ren, W., Yao, T., Xie, S., and He, Y.: Controls on the stable isotopes in precipitation
529 and surface waters across the southeastern Tibetan Plateau, *Journal of Hydrology*,
530 545, 276–287, <https://doi.org/10.1016/j.jhydrol.2016.12.034>, 2017.

531 Rindsberger, M., Magaritz, M., Carmi, I., and Gilad, D.: The relation between air mas
532 s trajectories and the water isotope composition of rain in the Mediterranean Sea
533 area, *Geophysical Research Letters*, 10, 43–46, [https://doi.org/10.1029/GL010i001p0](https://doi.org/10.1029/GL010i001p00043)
534 0043, 1983.

535 Risi, C., Noone, D., Worden, J., Frankenberg, C., Stiller, G., Kiefer, M., Funke, B., W
536 alker, K., Bernath, P., Schneider, M., Wunch, D., Sherlock, V., Deutscher, N., Grif
537 fith, D., Wennberg, P. O., Strong, K., Smale, D., Mahieu, E., Barthlott, S., Hase,
538 F., García, O., Notholt, J., Warneke, T., Toon, G., Sayres, D., Bony, S., Lee, J.,

539 Brown, D., Uemura, R., and Sturm, C.: Process-evaluation of tropospheric humidit
540 y simulated by general circulation models using water vapor isotopologues: 1. Co
541 mparison between models and observations, *Journal of Geophysical Research: Atm*
542 *ospheres*, 117, <https://doi.org/10.1029/2011JD016621>, 2012.

543 Ruan, J., Zhang, H., Cai, Z., Yang, X., and Yin, J.: Regional controls on daily to inte
544 rannual variations of precipitation isotope ratios in Southeast China: Implications f
545 or paleomonsoon reconstruction, *Earth Planet. Sci. Lett.*, 527, 115794, [https://doi.or](https://doi.org/10.1016/j.epsl.2019.115794)
546 [g/10.1016/j.epsl.2019.115794](https://doi.org/10.1016/j.epsl.2019.115794), 2019.

547 Scholl, M. A. and Murphy, S. F.: Precipitation isotopes link regional climate patterns t
548 o water supply in a tropical mountain forest, eastern Puerto Rico, *Water Resource*
549 *s Research*, 50, 4305–4322, <https://doi.org/10.1002/2013WR014413>, 2014.

550 Shi, M., Worden, J. R., Bailey, A., Noone, D., Risi, C., Fu, R., Worden, S., Herman,
551 R., Payne, V., Pagano, T., Bowman, K., Bloom, A. A., Saatchi, S., Liu, J., and F
552 isher, J. B.: Amazonian terrestrial water balance inferred from satellite-observed w
553 ater vapor isotopes, *Nat. Commun.*, 13, 2686, <https://doi.org/10.1038/s41467-022-30>
554 [317-4](https://doi.org/10.1038/s41467-022-30317-4), 2022.

555 Shi, Y., Wang, S., Wang, L., Zhang, M., Argiriou, A. A., Song, Y., and Lei, S.: Isotop
556 ic evidence in modern precipitation for the westerly meridional movement in Cent
557 ral Asia, *Atmospheric Res.*, 259, 105698, <https://doi.org/10.1016/j.atmosres.2021.105>
558 [698](https://doi.org/10.1016/j.atmosres.2021.105698), 2021.

559 Sun, C., Chen, Y., Li, J., Chen, W., and Li, X.: Stable isotope variations in precipitati
560 on in the northwesternmost Tibetan Plateau related to various meteorological contr
561 olling factors, *Atmospheric Res.*, 227, 66–78, <https://doi.org/10.1016/j.atmosres.2019.>
562 [04.026](https://doi.org/10.1016/j.atmosres.2019.04.026), 2019.

563 Sun, C., Tian, L., Shanahan, T. M., Partin, J. W., Gao, Y., Piatrunia, N., and Banner,
564 J.: Isotopic variability in tropical cyclone precipitation is controlled by Rayleigh d
565 istillation and cloud microphysics, *Commun Earth Environ*, 3, 1–10, [https://doi.org/](https://doi.org/10.1038/s43247-022-00381-1)
566 [10.1038/s43247-022-00381-1](https://doi.org/10.1038/s43247-022-00381-1), 2022.

567 Sun, Q., Miao, C., Duan, Q., Ashouri, H., Sorooshian, S., and Hsu, K.-L.: A Review
568 of Global Precipitation Data Sets: Data Sources, Estimation, and Intercomparisons,

569 Reviews of Geophysics, 56, 79–107, <https://doi.org/10.1002/2017RG000574>, 2018.

570 Tharammal, T., Bala, G., and Noone, D.: Impact of deep convection on the isotopic a
571 mount effect in tropical precipitation, *Journal of Geophysical Research: Atmospher*
572 *es*, 122, 1505–1523, <https://doi.org/10.1002/2016JD025555>, 2017.

573 V.I. Ferronsky and V.A. Polyakov, P. (Eds.): *Isotopes of the Earth's Hydrosphere*, Sprin
574 ger Publications, ISBN: 978-94-007-2855-4, 2012.

575 Wang, Q., Zhai, P.-M., and Qin, D.-H.: New perspectives on ‘warming–wetting’ trend
576 in Xinjiang, China, *Advances in Climate Change Research*, 11, 252–260, <https://doi.org/10.1016/j.accre.2020.09.004>, 2020.

577

578 Wang, S., Jiao, R., Zhang, M., Crawford, J., Hughes, C. E., and Chen, F.: Changes in
579 Below-Cloud Evaporation Affect Precipitation Isotopes During Five Decades of
580 Warming Across China, *Journal of Geophysical Research: Atmospheres*, 126, e202
581 0JD033075, <https://doi.org/10.1029/2020JD033075>, 2021.

582 Wang, S., Lei, S., Zhang, M., Hughes, C., Crawford, J., Liu, Z., and Qu, D.: Spatial
583 and Seasonal Isotope Variability in Precipitation across China: Monthly Isoscapes
584 Based on Regionalized Fuzzy Clustering, *J. Clim.*, 35, 3411–3425, <https://doi.org/10.1175/JCLI-D-21-0451.1>, 2022.

585

586 Wang, Y., Liu, X., and Herzschuh, U.: Asynchronous evolution of the Indian and East
587 Asian Summer Monsoon indicated by Holocene moisture patterns in monsoonal c
588 entral Asia, *Earth-Sci. Rev.*, 103, 135–153, <https://doi.org/10.1016/j.earscirev.2010.09.004>, 2010.

589

590 Yan, D., Xu, H., Lan, J., Zhou, K., Ye, Y., Zhang, J., An, Z., and Yeager, K. M.: Sol
591 ar activity and the westerlies dominate decadal hydroclimatic changes over arid C
592 entral Asia, *Global and Planetary Change*, 173, 53–60, <https://doi.org/10.1016/j.gloplacha.2018.12.006>, 2019.

593

594 Yao, J., Chen, Y., Chen, J., Zhao, Y., Tuoliewubieke, D., Li, J., Yang, L., and Mao,
595 W.: Intensification of extreme precipitation in arid Central Asia, *Journal of Hydrol*
596 *ogy*, 598, 125760, <https://doi.org/10.1016/j.jhydrol.2020.125760>, 2021.

597 Yao, T., Masson-Delmotte, V., Gao, J., Yu, W., Yang, X., Risi, C., Sturm, C., Werner,
598 M., Zhao, H., He, Y., Ren, W., Tian, L., Shi, C., and Hou, S.: A review of clim

599 atic controls on $\delta^{18}\text{O}$ in precipitation over the Tibetan Plateau: Observations and s
600 imulations, *Rev. Geophys.*, 51, 525–548, <https://doi.org/10.1002/rog.20023>, 2013.

601 Zhang, F., Huang, T., Man, W., Hu, H., Long, Y., Li, Z., and Pang, Z.: Contribution
602 of Recycled Moisture to Precipitation: A Modified D-Excess-Based Model, *Geoph*
603 *ysical Research Letters*, 48, e2021GL095909, <https://doi.org/10.1029/2021GL095909>,
604 2021a.

605 Zhang, J., Yu, W., Jing, Z., Lewis, S., Xu, B., Ma, Y., Wei, F., Luo, L., and Qu, D.:
606 Coupled Effects of Moisture Transport Pathway and Convection on Stable Isotop
607 es in Precipitation across the East Asian Monsoon Region: Implications for Paleoc
608 limate Reconstruction, *J. Clim.*, 1–41, <https://doi.org/10.1175/JCLI-D-21-0271.1>, 20
609 21b.

610 Zhang, Q., Gu, X., Singh, V. P., Sun, P., Chen, X., and Kong, D.: Magnitude, frequen
611 cy and timing of floods in the Tarim River basin, China: Changes, causes and im
612 plications, *Glob. Planet. Change*, 139, 44–55, [https://doi.org/10.1016/j.gloplacha.201](https://doi.org/10.1016/j.gloplacha.2015.10.005)
613 [5.10.005](https://doi.org/10.1016/j.gloplacha.2015.10.005), 2016.

614 Zhang, X., Lu, C., and Guan, Z.: Weakened cyclones, intensified anticyclones and rece
615 nt extreme cold winter weather events in Eurasia, *Environ. Res. Lett.*, 7, 044044,
616 <https://doi.org/10.1088/1748-9326/7/4/044044>, 2012.

617 Zhu, G., Guo, H., Qin, D., Pan, H., Zhang, Y., Jia, W., and Ma, X.: Contribution of
618 recycled moisture to precipitation in the monsoon marginal zone: Estimate based
619 on stable isotope data, *J. Hydrol.*, 569, 423–435, [https://doi.org/10.1016/j.jhydrol.20](https://doi.org/10.1016/j.jhydrol.2018.12.014)
620 [18.12.014](https://doi.org/10.1016/j.jhydrol.2018.12.014), 2019.

621 Zhu, G., Liu, Y., Shi, P., Jia, W., Zhou, J., Liu, Y., Ma, X., Pan, H., Zhang, Y., Zhan
622 g, Z., Sun, Z., Yong, L., and Zhao, K.: Stable water isotope monitoring network
623 of different water bodies in Shiyang River basin, a typical arid river in China, *E*
624 *arth Syst. Sci. Data*, 14, 3773–3789, <https://doi.org/10.5194/essd-14-3773-2022>, 20
625 2.

626 Zhu, G., Liu, Y., Wang, L., Sang, L., Zhao, K., Zhang, Z., Lin, X., and Qiu, D.: The
627 isotopes of precipitation have climate change signal in arid Central Asia, *Global*
628 *and Planetary Change*, 104103, <https://doi.org/10.1016/j.gloplacha.2023.104103>, 202

629 3.

630 Zhu, G., Zhang, Z., Guo, H., Zhang, Y., Yong, L., Wan, Q., Sun, Z., and Ma, H.: Be
631 low-Cloud Evaporation of Precipitation Isotopes over Mountains, Oases, and Deser
632 ts in Arid Areas, *J. Hydrometeorol.*, 22, 2533–2545, [https://doi.org/10.1175/JHM-D-](https://doi.org/10.1175/JHM-D-20-0170.1)
633 20-0170.1, 2021.

634