

A global dataset of the shape of drainage systems

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Abstract. Drainage basins delineate ~~the~~ Earth's land surface into individual water collection units. Basin shape and river sinuosity determine water and sediment dynamics, affecting landscape evolution and connectivity between ecosystems and freshwater species. However, a high-resolution global dataset for the boundaries and geometry of basins is still missing. Using a 90-meter resolution digital elevation model, we measured the areas, lengths, widths, aspect ratios, slopes, and elevations for basins ~~greater than~~over 50 km² globally. Additionally, we calculated the lengths and sinuosities of the longest river channels within these 0.67 million basins. We built a new global dataset, Basin90m, to present the basins and rivers, as well as their morphological metrics. To highlight the use cases of Basin90m, we explored ~~differences between the nine stream orders, spatial distribution of drainage systems, and the~~ correlations ~~between among~~ morphological metrics, such as Hack's law. By comparing with ~~HydroSHEDS, HydroSHEDS, HydroATLAS, and Google Earth images, and a few other datasets,~~ we ~~have~~ demonstrated the high accuracy of Basin90m. Basin90m, available in Shapefile format, can be used in various GIS platforms, including QGIS, ArcGIS, and GeoPandas. Basin90m has substantial application prospects in geomorphology, hydrology, and ecology.

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1 Introduction

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Drainage divides and rivers are among the most recognizable features on Earth's surface. The shape of drainage basins and rivers holds significant implications for landscape evolution processes and dynamics (Kirchner et al., 2001; Shelef, 2018; Ielpi et al., 2023). With equal basin lengths, a broader basin collects more precipitation thus offering a more stable discharge and water level, reducing the risk of river drying up. A stable water level benefits ecological integrity, water use, and navigation (Datry et al., 2023)~~delivers more water and sediments downstream.~~ Broader basins typically have longer and more intricate river networks, affecting flood risk associated with both extreme rainfall and glacial lake outburst events. In broader basins, rainwater takes more time to reach the main stream, indicating a longer arrival time of flood peaks. A glacial lake outburst flood travels a longer path in a broader basin, losing greater

40 energy and surface runoff. Hence, communities downstream of a broader high-mountain catchment encounter less threat from a single glacial lake outburst flooding. However, as broader basins have more tributaries with risks of glacial lake outbursts, timely and accurate flood early warnings require deploying more seismic stations within the basin (Maurer et al., 2020; Cook et al., 2021).

45 Furthermore, broader basins can contain a greater variety of tributaries and habitat types. This heterogeneity of habitats allows these basins to host a greater diversity of species, especially aquatic, amphibious, and riparian organisms, by providing more ecological niches across the landscape (Matthews et al., 1998). Similarly, meandering rivers create a mosaic of habitats with varying flow velocities, depths, and substrates, supporting the diversity of aquatic organisms (Nagayama and Nakamura, 2017; Rhoads et al., 2003; Yu et al., 2022).

50 Meanwhile, given the same elevation drop, a meandering river has a milder channel gradient than a straight river and therefore ~~may feature~~ smaller-lower erosion rates. ~~Meanwhile, meandering rivers create a mosaic of habitats with varying flow velocities, depths, and substrates, supporting the diversity of aquatic organisms (Nagayama and Nakamura, 2017; Rhoads et al., 2003; Yu et al., 2022).~~ Additionally, the shape of drainage systems has been argued to be related to climatic and tectonic conditions (Castelltort et al., 2012; Ielpi et al., 2023; Luo et al., 2023; Sreedevi et al., 2009; Strong and Mudd, 2022), and could therefore be used as an archive to reconstruct Earth's history. Accordingly, a global dataset on the shape of drainage systems benefits scientists and policymakers in geomorphology, hydrology, and ecology, fostering interdisciplinary collaborations.

60 Since the 21st century~~In the past three decades~~, digital elevation models (DEMs) have been utilized to produce several global-scale drainage basin and river databases (Allen and Pavelsky, 2018; Amatulli et al., 2022; Lin et al., 2021; Masutomi et al., 2009; Shen et al., 2017; Vörösmarty et al., 2000). Using 1-km resolution DEM, the U.S. Geological Survey (USGS) developed HYDRO1k (USGS, 2000), a global hydrological dataset providing vector basin boundaries and river channels. HydroSHEDS (Lehner and Grill, 2013; Lehner et al., 2008) provides 500-m resolution global basins and rivers and their basic metrics, such as area, river length, and stream order. ~~HydroSHEDS encompasses one million drainage basins.~~ HydroSHEDS basin boundaries and a 500-m resolution DEM were utilized to calculate morphometric indices for 26272 drainage basins worldwide with an area larger than 100 km² (Guth, 2011). These indices include basin elevation and slope. HydroATLAS (Linke et al., 2019) was also developed based on HydroSHEDS and incorporates 56 hydro-environmental attributes, including basin-mean slope and elevation (Linke et al., 2019).

75 The above datasets lack measurements of drainage basin length and aspect ratio. Shen et al. (2017) used 1-km DEM to obtain the global distribution of basin length and elongation ratio. However, the dataset is in raster format, without vectorial basin boundaries and rivers. More

80 ~~importantly, the spatial resolutions of all the above databases are relatively low. Recently,~~
~~w~~With the advancements in computer performance and algorithms, the 90-m resolution DEMs
are being used to establish global databases of drainage systems. For example, ~~the USGS~~
~~released-HDMA released by USGS, which has a resolution of 90 m (Verdin, 2017).~~ It includes
nearly 295000 drainage basins ~~but only contains information on drainage areas (Verdin, 2017)~~
~~and contains area attributes.~~ GRNWRZ (Yan et al., 2022) comprises a global river database at
a resolution of 90 m ~~that,~~ ~~includes~~ing information on ~~river the lengths of rivers (Yan et al.,~~
2022). This database offers the boundaries and areas of water resource zones, ~~which are~~ distinct
85 from drainage basins.

In summary, ~~catalogs of~~ drainage basins and rivers ~~are available~~~~have been mapped,~~ along
with measurements of area, slope, and elevation of basins. Yet, ~~many datasets were based on~~
~~DEMs with resolutions of 500 m or coarser, and~~ few works have focused on more complex
basin characteristics, such as aspect ratio, which describes the shape of drainage basins, and
90 sinuosity that characterizes the shape of river channels. ~~Furthermore, some databases only offer~~
~~raster formats without vector accessibility. Moreover, the download links of some datasets are~~
~~invalid (e.g., Vörösmarty et al., 2000; Guth, 2011).~~

Here, ~~we provide an updated global catalog of drainage systems, making all relevant files~~
~~and script available, and including a wide range of geometric characteristics.~~ ~~we~~We used a
95 global 90-m resolution DEM and obtained over 665000 drainage basins with a size ~~greater~~
~~than~~over 50 km². For each basin, we extracted the longest river channel that extends from
drainage divide to river mouth. Additionally, we measured parameters for each drainage system,
including stream order, the length, width, aspect ratio, slope, and elevation of basins, and the
length and sinuosity of rivers. The spatial distribution of drainage systems and their
100 morphological parameters ~~constitute form an drainage system shape dataset, improved~~
~~resolution global dataset,~~ Basin90m.

2 Methodology

105 ~~First, w~~We divided the global DEM into 130 segments to accommodate the computational
capabilities (Figs. 1 & 2a). ~~Then, w~~We selected a basin in Madagascar to demonstrate the ~~steps~~
~~processes~~ of obtaining drainage basins and their longest rivers from a DEM ~~through a series of~~
~~hydrological steps. These steps include~~First, ~~we d~~ ~~calculating~~ed flow direction and
accumulation for each point within the basin. Based on flow direction and accumulation, we
delineated basins and rivers of different stream orders (Fig. 2b). Once the spatial distribution of
110 basins and rivers was obtained, we measured parameters describing the size and shape of the
drainage systems (Fig. 2c). Finally, basins with over half of their area located in lakes or sandy
deserts were removed (Fig. 2d).

Conventional GIS platforms such as QGIS and ArcGIS provide various hydrology tools, such as extracting drainage basins and river channels. However, ~~in Basin90m,~~ we need to process 130 DEMs and calculate some metrics, such as basin width, ~~for which no tools are readily available within these GIS platforms that are not conveniently calculated within GIS software.~~ Therefore, we used the TopoToolbox software (Schwanghart and Scherler, 2014) to automate the extraction of drainage system and the calculation of various metrics in a single script. TopoToolbox is a Matlab toolbox for analyzing and manipulating geospatial data, supporting drainage basin delineation and topographic property calculation. In addition to utilizing the functions in TopoToolbox, we developed new algorithms to measure basin length and width. We integrated all these functions into an automated workflow that delineates drainage basins and channels across all stream orders and measures their morphological attributes.

All the operations described below in ~~Sections~~ 2.2 to 2.6 were automated within our TopoToolbox script. The script inputs a single DEM and outputs vector files (ESRI shapefiles) for basins and rivers. Refer to ~~Code code~~ availability statement for the download link of this script. Shapefile is a data format storing geographic vector data, including geometry and attributes. Parameters describing the morphology of ~~the~~ drainage systems ~~are were~~ stored in the attribute table of ~~the~~ basin shapefiles. The computation of all ~~the~~ 130 DEMs was carried out on the High Performance Computing platform of German Research Centre for Geosciences.

2.1 DEM preprocessing

The 90-m resolution Shuttle Radar Topography Mission (SRTM) DEM is a widely used dataset for global geomorphological analyses (Farr et al., 2007). It covers the Earth's land surface between 60 degrees north and 56 degrees south latitude. Due to limitations in computer memory, it is a common practice to partition ~~the~~ global DEM into smaller segments, ~~especially when working with a high-resolution global DEM~~ (Amatulli et al., 2022). To ensure the integrity of each drainage basin and river channel, we partitioned the global DEM based on drainage divides provided by HydroBASINS (Lehner and Grill, 2013). We used the Clip Tool in ArcGIS (version 10.2) to cut the global DEM into 130 segments (Figs. 1 & 2a). By utilizing drainage divides corresponding to the largest basins in a ~~specific~~ region, we ensure that continents ~~are are~~ cropped in a way that avoids splitting any basins internally. Many islands were treated as a single DEM segment, such as New Zealand (Fig. 1).

~~Using standard flow routing algorithms can lead to artifacts and unrealistic drainage networks, especially in lakes and deserts. Drainage systems within lakes and deserts are likely unrealistic. However, we did not remove the DEM of lakes and sandy deserts to maintain the integrity of major river basins such as the Nile River that traverses lakes and deserts. Further~~

processing of those regions was carried after obtaining the global drainage systems, as detailed in Sect. 2.7.

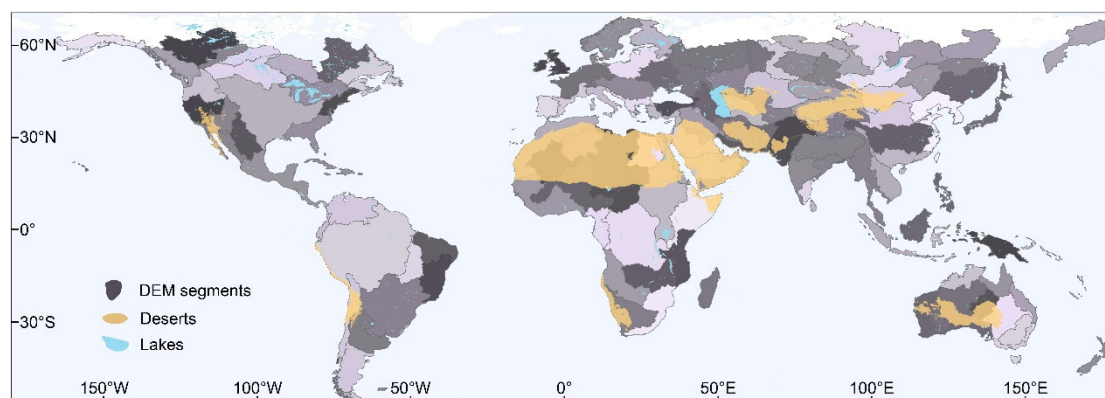


Figure 1. DEM segments (N=130) based on the drainage divides from HydroBASINS (Lehner and Grill, 2013). Different shades of gray ~~are were~~ used to distinguish different DEM segments. Lakes were obtained from HydroLAKES (Messenger et al., 2016). We consider regions with an aridity index less than 0.08 as sandy deserts ~~areas~~ (see Sect. ~~ion~~ 2.7 for details).

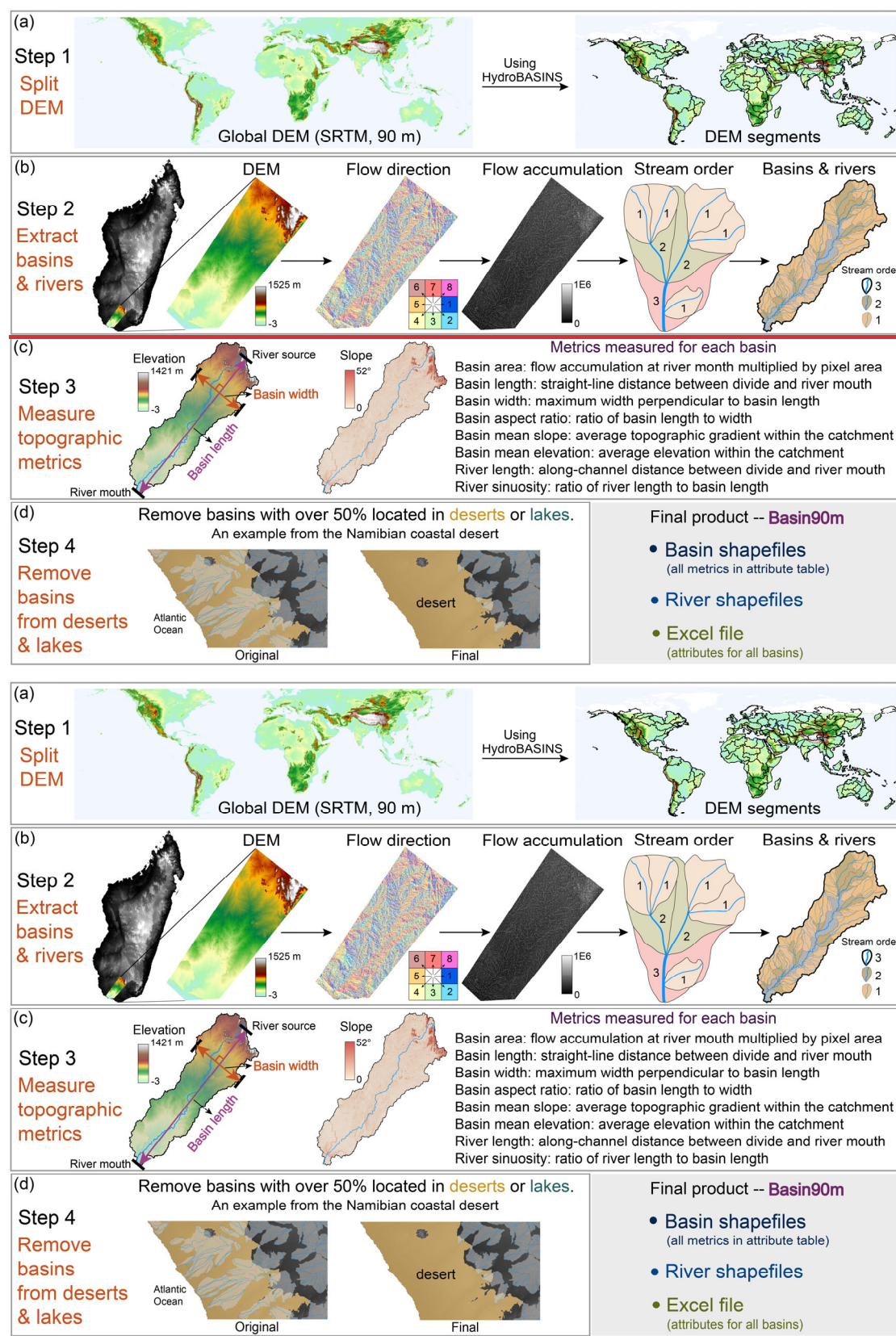
2.2 Flow direction

The first step to obtaining ~~drainage basins and river networks from a DEM is to calculate~~ drainage basins and river networks from a DEM is calculating flow direction and ~~flow~~ accumulation ~~for all cells within the DEM~~. Flow direction describes the drainage direction of each cell. Flow networks derived from DEMs are affected by measurement and data processing errors (Schwanghart et al., 2013). DEM elevations in valley bottoms can be overestimated due to steep hillslopes, water, and vegetation (Schwanghart and Scherler, 2017). Therefore, we used the carving method to generate flow directions to ensure channels are well connected despite local noise. ~~Carving is a process used to remove obstacles by ensuring that, while progressing downstream, no pixel has a higher elevation than its neighboring pixels upstream~~ (Lindsay, 2016; Schwanghart et al., 2013; Schwanghart and Scherler, 2017).

Once carving is complete, we used the D8 method to determine flow direction. ~~The D8 method is a commonly used approach for calculating river flow direction in hydrological analysis~~ (Tarboton, 1997). This method assumes that water flows directly downslope from each cell to one of its eight neighboring cells (Fig. 2b). ~~The D8 method is computationally efficient and suitable for high-resolution global hydrological analysis (Lehner and Grill, 2013; Yamazaki et al., 2019; Yan et al., 2022).~~ A single flow direction out of the eight possible directions is assigned to each cell of the DEM (Fig. 2b), by comparing the elevations of neighboring cells and identifying the steepest descent path. ~~This method assumes that water flows directly downslope from each cell to one of its eight neighboring cells (Fig. 2b).~~ ~~In the D8 method, despite various ways of numbering the eight flow directions, w~~ We adopt integers from 1 to 8

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to differentiate the possible directions. For example, if a cell has a flow direction of 5, it indicates that all water passing through that cell will be directed towards its left neighbor (Fig. 2b). Accordingly, we used D8 method to calculate flow direction for all the 130 DEMs.



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Figure 2. Flowchart illustrating the processes of extracting the global drainage systems and their

morphological metrics. (a) The global DEM was partitioned into 130 segments using drainage divides provided by HydroBASINS (Lehner ~~and Grill, 2013~~~~et al., 2008~~). (b) Menarandra River ~~Basin~~ basin in southern Madagascar was chosen as an example to demonstrate the steps of extracting drainage basins and the longest river channels with various stream orders. (c) Based on the spatial distribution of basins and river channels ~~obtained in the previous step~~, eight parameters describing the drainage system's size and morphology were automatically measured using a Matlab script. (d) Drainage basins that have over half of their area within lakes or sandy deserts were removed.

2.3 Flow accumulation

Flow direction is the basis for calculating flow accumulation. ~~Flow accumulation is~~ a measure ~~used in hydrological analysis to that~~ quantifies the cumulative number of cells contributing to a given point (Fig. 2b). ~~The flow is traced along the flow direction pathway starting from the upstream cells. Thus, flow accumulation starts from zero at the basin boundary, increasing downstream and reaches~~ a maximum at the river mouth. ~~The river~~ river mouths can be ~~a location where the river confluences~~ river meets a bigger river, a lakes, or the ~~ocean~~ seas. Therefore, ~~basin area can be obtained by multiplying flow accumulation at river mouth by the area of one pixel~~ multiplying the flow accumulation of the river mouth by the pixel area of the DEM equals basin area. With flow direction and accumulation, we can delineate stream orders and extract basins and rivers.

2.4 Stream order

Large drainage basins contain small catchments. We assigned Strahler stream orders to basins and their longest rivers for hierarchical classification to capture ~~this the~~ topological relationship between catchments. Strahler stream order ~~is a method used to classify and~~ quantifies the hierarchy of ~~basins and~~ river segments within a river network (Strahler, 1957). It assigns an ~~integer numerical value~~ to each segment based on the contributing tributaries (Fig. 2b). ~~First-order rivers are often found near the basin boundary. Stream order gradually increases downstream.~~ When two segments of the same order converge, they merge to form a new segment with an order ~~increased by one~~ higher. If two segments of different orders merge, the resulting segment inherits the higher order. For example, a first-order segment merging with a second-order segment results in a second-order segment. ~~First-order rivers are often found near the basin boundary. Stream order gradually increases downstream.~~

~~It is important to note that~~ The magnitude of stream order within a drainage basin depends on the area threshold of the first-order basin. The larger the threshold, the fewer orders in the entire catchment. Limited by computing resources, we only extracted ~~and analyzed~~ drainage basins with an area ~~≥ greater than or equal to~~ 50 km². ~~A drainage basin with a~~ An area of 50 km² contains about 6000 ~~elevation points~~ cells, sufficient to organize a drainage system. ~~In Basin90m, b~~ Basins with an areas of less than 50 km² were not extracted, thus not contributing to the stream order of ~~its~~ downstream basins. ~~However, those small basins contribute to water discharge of their downstream channels.~~

2.5 Drainage basins and the longest rivers

~~In TopoToolbox, basins and rivers are derived based on flow direction and accumulation.~~ We extracted only the longest river of each basin (Fig. 2b). In ~~natural landscapes~~, rivers commonly ~~begin to~~ develop downstream of ~~the~~ drainage divide by a certain distance, known as hillslope length. For example, the average hillslope lengths in Taiwan and Sicily are 1556 and 1756 m, respectively (He et al., 2021a). A common approach is to designate the point where the upstream area exceeds a channelization threshold ~~certain value~~ as the river source. ~~This value is termed as the channelization threshold.~~ For instance, the two latest global river databases, MERIT Hydro (Lin et al., 2021) and Hydrography90m (Amatulli et al., 2022), employ channelization thresholds of 1 and 0.05 km², respectively. ~~Because~~ Since Basin90m utilizes the straight-line distance between river source and mouth as ~~the~~ basin length, we selected a channelization threshold of zero, indicating that ~~the rivers in Basin90m~~ originate from drainage divides.

We use Menarandra River basin in Madagascar as an example to illustrate ~~our~~ the steps methodology for obtaining ~~catchments with area larger than 50 km², drainage systems using Menarandra River basin located in southern Madagascar as an example~~ (Fig. 2b,c). This ~~drainage~~ basin consists of three stream orders, including one order 3, ten order 2, and forty-six order 1 sub-basins. The third-order basin has an area of 8701 km² and a basin aspect ratio of 3.3. The corresponding river length and sinuosity are 295 km and 1.6, respectively.

2.6 ~~Measuring m~~Morphological indices

~~After obtaining the basin boundary and the longest river, we measured their basic geometric parameters~~ We measured the basic geometric parameters after obtaining basin boundaries and the longest rivers. For basins, the parameters include area, length, width, aspect ratio, and average slope and elevation (Fig. 2c). For rivers, we measured along-channel length and sinuosity (Fig. 2c). All parameters were automatically computed ~~in~~ with our TopoToolbox script. For instance, the straight-line distance between ~~the~~ river source (drainage divide) and ~~river~~ mouth of the longest channel is ~~used to describe~~ termed basin length. Basin width is defined as the greatest distance measured along a straight line perpendicular to the direction of basin length at the two points where the line intersects basin boundary. The sum of elevations of all pixels within a basin divided by the number of pixels yields the basin mean elevation. We report all ~~of these eight~~ metrics and stream order values for catchments ~~with a size larger than~~ over 50 km². The value of these metrics is stored in the attribute table of each basin shapefile data. See ~~data~~ data availability statement for details.

2.7 Removing basins from lakes and deserts

After obtaining the global distribution of drainage systems, we removed basins and rivers associated with lakes and sandy deserts (Fig. 2d). Basin90m aims to provide drainage systems created by surface water flow processes. While lakes are related to water flow, it is inappropriate to consider a drainage basin completely within a lake. We need a global lake database to remove drainage systems within lakes. Several global lake datasets are available, such as HydroLAKES (Messager et al., 2016), GloLakes (Hou et al., 2022), and Lake-TopoCat (Sikder et al., 2023). Here, we used HydroLAKES, which includes 1.4 million lakes worldwide, with an area range of 0.1-377002 km². Since the minimum basin area in Basin90m is 50 km², we only retained 3402 lakes with a size ~~greater than~~over 50 km² (Fig. 1).

Due to high permeability and intense evaporation, sandy deserts are typically unable to sustain surface water. Therefore, basins and rivers derived from DEMs within sandy deserts are unlikely to accurately reflect river dynamics. To our knowledge, no published global dataset for sandy deserts exists. As a result, we relied on using ~~the~~aridity index to determine the location of such regions. ~~The a~~Aridity index is a measure that quantifies the dryness or aridity of a region based on the ratio of precipitation to potential evapotranspiration. A lower aridity index indicates a drier environment. Therefore, aridity index is closely associated with desert formation (Gamo et al., 2013).-

Here, we used the Global-AI_PET_v3 (Zomer et al., 2022), a global aridity index dataset, to delineate sandy deserts. By comparing the regions obtained using multiple aridity index thresholds with the distribution of sandy deserts on Google Earth, we classified areas with aridity index values less than 0.08 as sandy deserts. This threshold captures the majority of sandy deserts while avoiding misclassifying excessive non-desert regions.-

Using the Python-based GeoPandas library, we removed basins ~~where over half of their area is~~ith over half of their area within lakes or sandy deserts (Fig. 2d). We also removed the river channels within those deleted basins. In addition ~~to filtering basins associated with lakes and deserts~~, we manually removed all drainage systems intersecting with the 60°N latitude line in ArcGIS to avoid incomplete river networks due to DEM coverage. In the end, out of the 840000 basins obtained in Section 2.5, approximately 170000 were removed, retaining 667629 basins and their longest rivers.

3 Results and discussions

Through the steps above, we obtained the spatial distribution of global drainage systems and their morphological metrics. Next, we analyze the distribution and correlations of these metrics in Basin90m. First, we present the differences in drainage systems with different stream orders. Then, we show the distribution of metrics that describe the morphology of ~~the~~ drainage systems

~~globally~~. Afterward, we present the interrelationships among these metrics. For example, utilizing Basin90m, we derived Hack's law (Hack, 1957) fitted from global data.

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3.1 Basins with different stream orders

Fig. 3a shows basins in Basin90m with stream orders ranging from 4 to 9. ~~The~~ ~~three~~ ninth-order basins are the Amazon, Nile, and Congo, with an area of 6038417, 2895947, and 3716644 km², respectively. Eighth-order basins are widely distributed globally, including the Mississippi in North America, the Río de la Plata in South America, the Orange in Africa, the Yangtze and Mekong in Asia, and the Murray-Darling in Australasia. Basins with the highest stream order in a region often occupy a significant portion of the area. For example, Madagascar has a basin with a stream order 6, covering an area of 59200 km², approximately 10% of Madagascar's total area. Taiwan island has two basins with a stream order of 4, each with an area of about

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30000 km², together occupying one-sixth of the island.

Fig. 3b-g illustrates the statistics of drainage systems with different stream orders. As stream order increases, the number of basins and total river length exponentially decrease (Fig. 3b,c). The number of first-order basins is 521857, accounting for 78% of the total basin count. There are 661397 basins with orders 1-3, representing 99% of the total basin count. In contrast,

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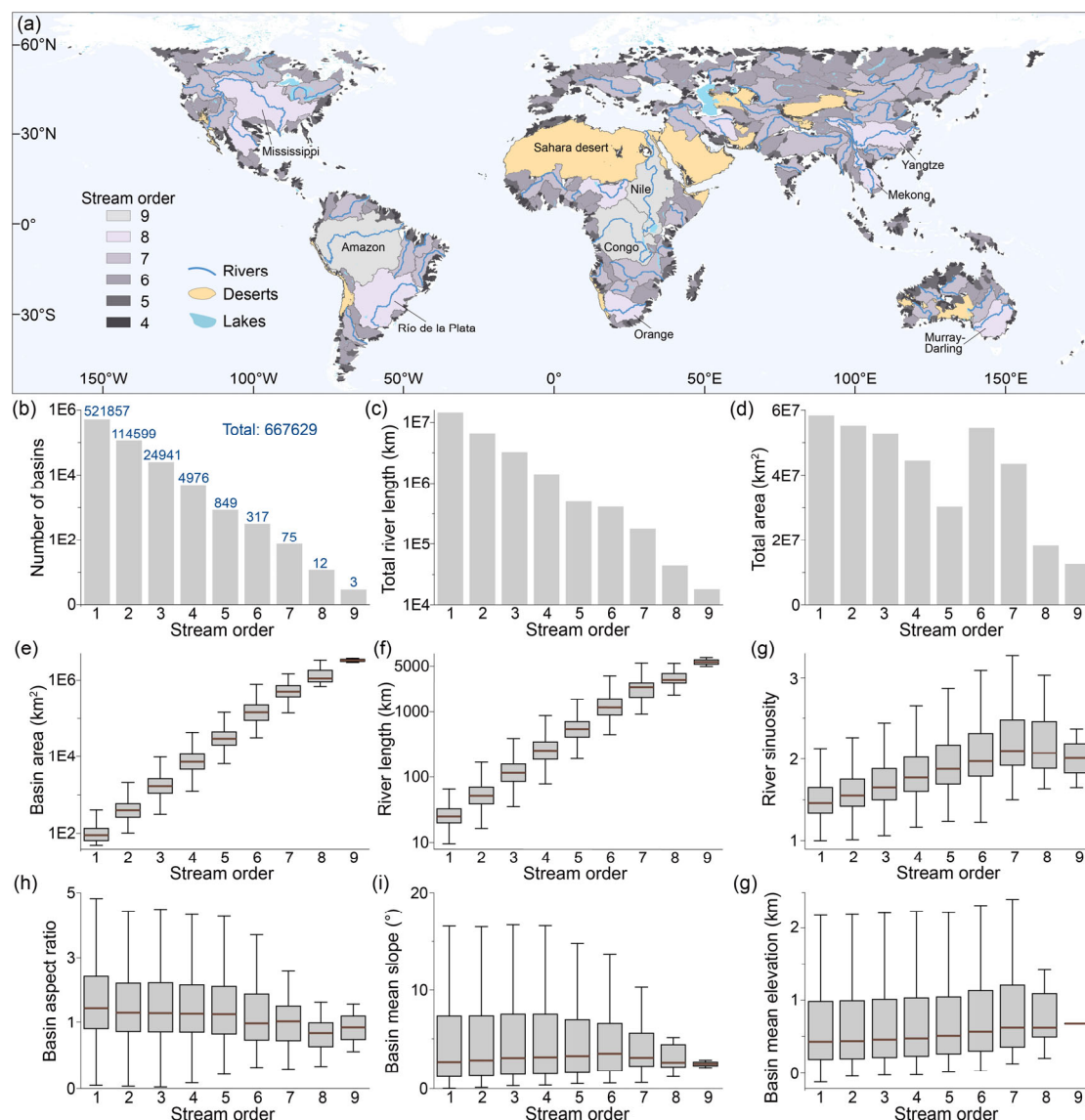
only 90 basins are of the highest stream orders (7-9). The total basin area is determined by multiplying the basin count with the average basin area. As stream order increases, the average basin area increases while the basin count decreases. Therefore, total basin area does not ~~monotonically systematically~~ change with stream order (Fig. 3d). The average basin area and river length increase with increasing stream order because higher-order basins encompass

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lower-order basins (Fig. 3e,f).-

River sinuosity generally increases with ~~higher~~ stream orders (Fig. 3g), as higher-order basins provide more space for rivers to meander (Biron et al., 2014). The lack of significant changes in basin aspect ratio with increasing stream order (Fig. 3h), along with the absence of changes in basin aspect ratio with increasing stream order (Fig. 3h), along with the absence of a correlation between basin topography (slope and elevation) and stream order (Fig. 3i,g), supports the concept of self-similarity in drainage basins (Bennett and Liu, 2016; Mantilla et al., 2010; Sassolas-Serrayet et al., 2018).

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330 **Figure 3.** Global drainage systems with nine stream orders. (a) The spatial distribution of basins
 with orders from four to nine. Rivers with stream orders from 7 to 9 are displaced. (b-g)
 Morphological metrics displayed by stream orders.

3.2 Distribution of morphological metrics

335 We present the probability density ~~iesy distributions~~ and spatial patterns of metrics that describe
 the ~~morphological morphology characteristics~~ of ~~the global~~ drainage systems. Most
 morphological metrics have left-skewed log-normal distributions (Fig. 4). Ninety-five percent
 of basins have an area smaller than 1000 km². There are only 2868 basins with a size larger than
 10000 km² (Fig. 4a). Most (88%) basins have a length ranging from 10 to 50 km, with an
 average value of 25 km (Fig. 4b). The number of basins with lengths exceeding 100 km is 11357,
 340 accounting for only 1.7%. The average basin length is roughly twice the ~~average~~ basin width.

Many basins (N = 89673) have an average width ranging from 5 to 25 km, constituting
 87% of the total (Fig. 4c). The number of basins with aspect ratio between one and five is

598426, accounting for 90% of all basins (Fig. 4d). A basin aspect ratio less than one typically indicates a ~~significant~~ deviation between the overall flow direction and ~~basin the~~ elongation direction ~~of the basin~~. ~~This scenario is observed in only~~ ~~These are only~~ 5% of the basins (N = 34357), suggesting ~~that a deviation between the elongation and flow direction in river basins is~~ ~~rare~~ a rare deviation between basin elongation and overall water flow.

The average slope shows a wider distribution compared to the other parameters. It is distributed between 0 and 40°, with an average value of 5.4° (Fig. 4e). ~~The distribution of Basin~~ ~~basin~~ mean elevation ~~directly reflects the distribution of elevations globally with~~ ~~has~~ multiple peaks (Fig. 4f). Among them, the peak at an elevation of 4500 m corresponds to ~~the~~ Tibetan Plateau and Andes ~~(Fig. 4f)~~. Over 95% of the river lengths are smaller than 100 km (Fig. 4g). The number of basins with river lengths exceeding 100 km and 1000 km is 31392 and 370, respectively. Since river sinuosity is the ratio between the along-channel ~~length~~ and ~~the~~ straight-line distance from ~~the source to the~~ drainage divide to river mouth ~~of a river~~, the minimum ~~sinuosity value~~ is one. Basins with a river sinuosity less than two account for 91%, with an average value of 1.6 (Fig. 4h).

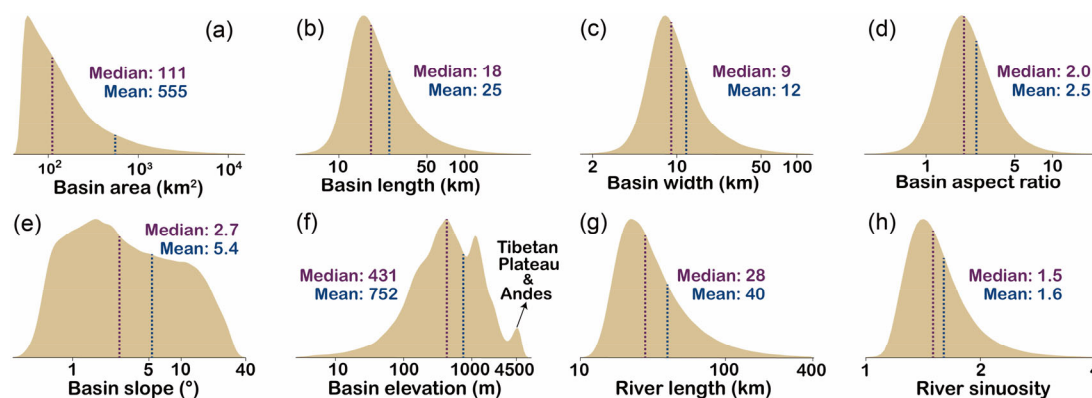
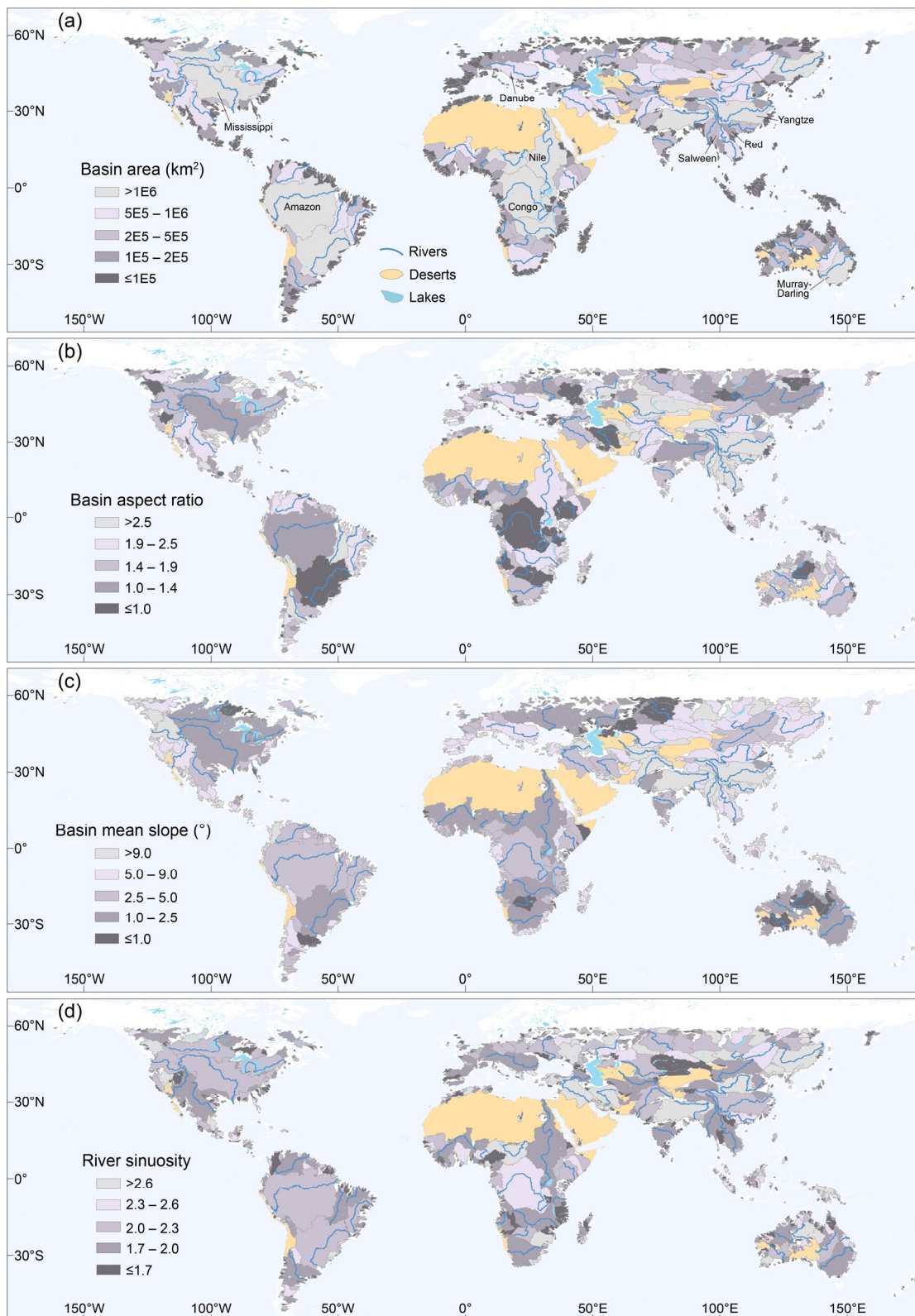


Figure 4. The probability density estimation shows the distribution of morphological metrics. All basins in Basin90m were used for probability density estimation without distinguishing stream orders. The ~~X~~_Y-axis of all subfigures is logarithmically scaled with a base of ten. (a) Basin area. (b) Basin length. (c) Basin width. (d) Basin aspect ratio. (e) Basin mean topographic slope. (f) Basin mean elevation. (g) ~~River~~ ~~The~~ length ~~of the longest river channels~~. (h) ~~River~~ ~~The~~ sinuosity ~~of the longest river channels~~. Purple and blue dashed lines indicate the median and the mean, respectively.

Basin90m includes eight metrics for each drainage system (Figs. 2c & 4). Here, we present the spatial distribution of four ~~metrics of them~~ (Fig. 5). The first parameter is drainage area, ~~which describes~~ ~~ing~~ basin size. There are 12 basins with an area larger than ~~1~~ ~~one~~ million km², with stream orders ranging from 7 to 9 (Fig. 5a). The second parameter is aspect ratio, illustrating basin shape. A higher aspect ratio indicates a more elongated basin. There is no clear relationship between aspect ratio and stream order (Fig. 3h). As a result, the most elongated basins can be ~~both~~ large and small (Fig. 5b). Numerous elongated basins characterize the ~~periphery of~~ Tibetan Plateau ~~and its surrounding areas~~. An aspect ratio smaller than one indicates that the overall flow direction ~~of the river~~ is not aligned with ~~the~~ ~~basin~~ elongation

direction ~~of the basin~~. For example, the Congo River flows from west to east, while the elongation ~~longest~~ direction of the basin is north-south. This results in an aspect ratio of 0.9 for the Congo Basin.

380 The third metric is the average slope of the basin, demonstrating the topographic variations within the catchment. Due to the significant elevation differences between ~~the~~ Tibetan Plateau and the surrounding plains, steep basins ~~have~~ developed around the plateau (Fig. 5c). The average slopes of the Yangtze, Salween, and Red Basins around the plateau margin are 13, 17, and 18°, respectively. In contrast, the slopes of the Murray-Darling, Mississippi, and Amazon ~~basins~~ Basins are only 1, 2, and 3°, respectively. The fourth parameter is river sinuosity, 385 describing the shape of the longest river. A high sinuosity indicates a meandering river, while a sinuosity closer to one indicates a straight river. The sinuosity of the Nile, Danube, Amazon, and Mississippi rivers are 1.6, 1.8, 2, and 2.3, respectively (Fig. 5d).



390 **Figure 5.** The global distribution of drainage systems with various metrics. Note that higher-order basins cover the lower-order basins. Drainage basins colored by basin area (a), basin aspect ratio
 | (b), basin mean topographic slope (c), and river sinuosity (d). Rivers (blue) with stream orders from 7 to 9 are displaced.

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3.3 Relationship between morphological metrics

The eight metrics in Basin90m describe the size and shape of drainage systems from various perspectives. The correlation among these parameters is crucial for understanding landscape dynamics, thus we present a visualization of their relationships (Fig. 6). Overall, there is a strong correlation among parameters that describe similar features (Fig. 6a). For instance, the correlation coefficient between basin length and river length is 0.97, suggesting that river channels and basins tend to grow or shrink simultaneously. The expansion and contraction of basins are achieved through drainage divide migration, which redistributes the lengths of river channels on both sides of the drainage divide (Habousha et al., 2023; He et al., 2021b). Conversely, parameters describing different features typically exhibit a weak correlation. For example, the correlation coefficient between the area representing basin size, and the slope describing topography, is approximately zero (Fig. 6a).

Fig. 6b displays 2D density plots of correlated parameters. For example, Basin area increases with increasing basin length and width. This result is expected since basins often become wider as they grow in length (correlation coefficient of 0.83, Fig. 6a), and both length and width determine basin area. Another example is the relatively high correlation between elevation and slope, which indicates that basins with steeper slopes tend to develop at higher elevations. It is important to note that the correlations shown in Fig. 6 represent the overall trend, with numerous exceptions exist. For instance, high-altitude plateaus have low slopes, while areas with only a few hundred meters can exhibit high slopes due to deep-cut gorges.

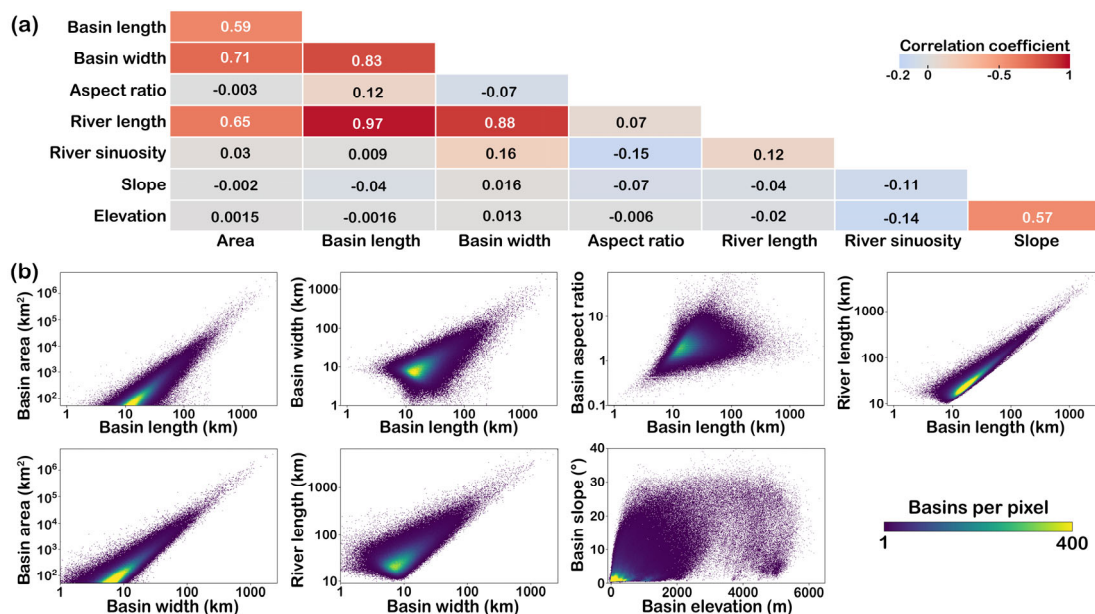
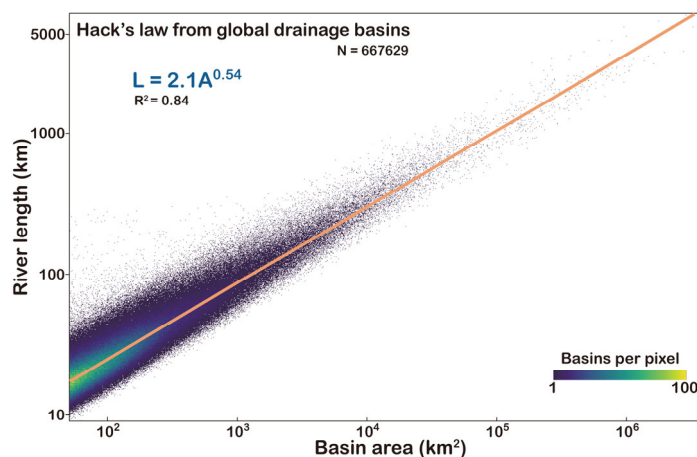


Figure 6. The correlation among metrics. (a) The heatmap shows the linear correlation coefficients among eight metrics, highlighted by the heatmap. The correlation coefficient indicates the extent and direction of the linear relationship between two variables. It ranges from -1 to 1, where positive values signify a positive correlation, with a higher absolute value indicating a stronger relationship. (b) 2D density plot revealing relationships between selected metrics.

425 3.4 Hack's law

Hack's law is an empirical relationship that relates river length to basin area (Hack, 1957). It can be expressed as $L = kA^h$, where L represents the length of the longest river measured from drainage divide to river mouth, A is drainage area, k and h are constant and exponent, respectively. Hack's law is a fundamental concept in geomorphology and hydrology, ~~and is~~ crucial for understanding ~~the dynamics of~~ river networks dynamics. In addition, Hack's law derived from Earth can potentially be compared to valley systems on Mars, providing insights into its climates and hydrological processes (Luo et al., 2023; Penido et al., 2013; Som et al., 2009). Previous studies have obtained different values for k and h in Hack's law, which vary across regions. The k ranges from one to five, while the range for ~~the~~ h is typically 0.4-0.7 (Hack, 1957; Luo et al., 2023; Mueller, 1972; O'Malley, 2020; Sassolas-Serrayet et al., 2018; Yi et al., 2018). Recently, based on 3685 ~~catchments globally~~ global catchments, Hack's law with an exponent of 0.56 was obtained (O'Malley, 2020).-

Here, we utilized ~~the~~ basin areas and river lengths in Basin90m to establish a new global Hack's law with $k = 2.1 \text{ km}^{-0.08}$ and $h = 0.54$ (Fig. 7). These values not only match regional case studies (Hack, 1957; He et al., 2021a; Montgomery and Dietrich, 1992; Sassolas-Serrayet et al., 2018; Yi et al., 2018; Yuan et al., 2023), but also align with ~~earlier~~ results derived from global drainage systems (O'Malley, 2020).



445 **Figure 7.** Hack's law fitted based on ~~the~~ basin area (A) and river length (L). We used all basins in Basin90m to fit Hack's law, including all stream orders.

4 Validations and limitations

Due to the absence of a published database for a direct comparison with the eight parameters contained in Basin90m, we used the Hydrology Tool in ArcGIS (~~version 10.2~~) to extract the drainage divide and main channel of the Moche ~~Basin~~ basin in Peru. ~~Here, we used the same DEM as Basin90m, which is the 90-m resolution SRTM DEM (Farr et al., 2007).~~ On this foundation, we validated the accuracy of the spatial position of basin ~~boundaries~~ boundary and

the longest river channel against HydroSHEDS (Lehner and Grill, 2013), the drainage system extracted by ArcGIS, and ~~the~~ Google Earth images. We then compared the morphological metrics in Basin90m against HydroATLAS (Linke et al., 2019) and the measurements from ArcGIS. Furthermore, we selected ten representative drainage basins spanning the North American continent to compare drainage areas with HydroATLAS. In addition, we discussed the limitations of Basin90m.

4.1 Spatial accuracy of drainage system

The accuracy of basin boundaries and river channels directly affects the calculation of the sizes and shapes of drainage systems. We used the Moche River basin in Peru as an example to verify the spatial accuracy of Basin90m, because it contains diverse terrain features (Fig. 8a). Additionally, the presence of the Moche Civilization in the downstream plain of the Moche River indicates that human modification on the landscape has been ongoing for approximately three thousand years (Toyne et al., 2014). ~~The~~ Moche River originates in the Andes and flows into the Pacific Ocean. Moche River basin encompasses an area of 2143 km², with its highest altitude reaching 4257 m. Nearly half of the drainage basin is located in the upstream low-relief plateau. The middle reaches consist of a deep canyon that occupies 40% of the area. The downstream area is a ~~flat~~ plain covering 10% of the ~~drainage~~ basin, and the river only decreases in elevation by 150 m over 20 km along the plain.

We compared the spatial accuracy of drainage systems in Basin90m with two datasets. The first one is HydroSHEDS which consists of HydroRIVERS and HydroBASINS (Lehner and Grill, 2013). HydroSHEDS has a spatial resolution of 500 m and has been widely used in various fields such as geomorphology, hydrology, ecology, climatology, and geohazards (McEwan et al., 2023; Palmer et al., 2023; Tu et al., 2023). The second dataset we compare with Basin90m is the drainage system of Moche ~~basin-basin~~ we extracted from ArcGIS using the same 90-m SRTM DEM (Farr et al., 2007). Due to the different DEM resolutions and methods in the three data, the basins and rivers only partially overlap (Fig. 8a).

In the deep canyon area (Fig. 8b), HydroRIVERS is coarse and does not match the river channel. In comparison, although Basin90m does not fully align with the ~~real-actual~~ river channel, there is a noticeable improvement in accuracy. The accuracy of the river extracted using ArcGIS falls between Basin90m and HydroRIVERS. The differences between the three are more pronounced in the ~~flat plain-of the downstream area~~ (Fig. 8c). HydroRIVERS lies along the floodplain rather than the river channel. In contrast, even in such a low-relief area with human modifications, Basin90m follows the position of the river channel. ~~Again, t~~The performance of the river obtained using ArcGIS falls between the two datasets.

We selected a segment of the drainage divide to compare the accuracy of basin boundary

(Fig. 8d). The drainage divide in Basin90m follows the ridge. However, HydroBASINS does not follow the ridge and even cuts through a deeply incised river channel. Although the drainage divide extracted using ArcGIS generally follows the ridge ~~with some degree of accuracy~~, it is not as accurate as Basin90m. In summary, ~~in the Moche basin, Basin90m exhibits a higher spatial accuracy~~ Basin90m exhibits a higher spatial accuracy in the Moche basin than HydroSHEDS and drainage systems extracted using ArcGIS.

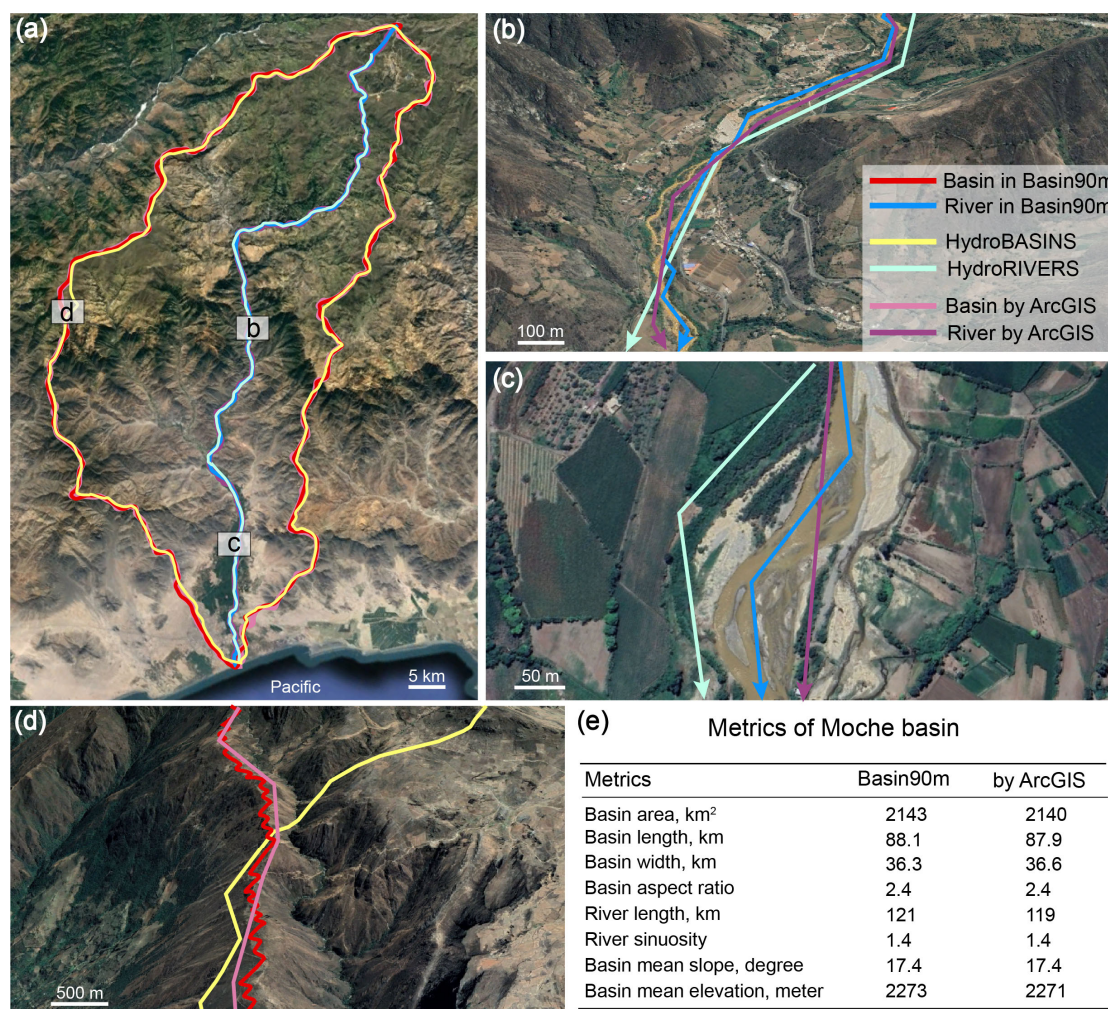


Figure 8. ~~The s~~patial accuracy of the drainage system in Basin90m, compared against HydroSHEDS (Lehner and Grill, 2013), drainage system extracted by ArcGIS, and Google Earth images. The Moche basin in Peru serves as an example that ~~combines a low-relief plateau, deep canyons, and farming land in plains, simultaneously~~ simultaneously combines a low-relief plateau, deep canyons, and farming land in plains. (a) The drainage system of the Moche basin. A comparison between river channels in deep canyon region (b) and flat region (c). (d) A comparison between drainage divides. (e) A comparison of the metrics for Moche ~~Basin-basin~~ Basin-basin between Basin90m and those extracted by ArcGIS.

4.2 Accuracy of morphological metrics

The morphological metrics that can be directly compared between the published databases and Basin90m only include basin area, slope, and elevation. For example, in HydroATLAS (Linke et al., 2019), the area, slope, and elevation of the Moche ~~Basin-basin~~ Basin-basin are 2125 km², 17°2267 m,

and ~~17°22'67 m~~, respectively. This means the difference between Basin90m and HydroATLAS in Moche ~~Basin-basin~~ is less than 1%. To validate the accuracy of the remaining five parameters included in Basin90m, we used the results of the Moche ~~Basin-basin~~ extracted by ArcGIS.

~~_____~~ The primary difference in the extraction of drainage system between ArcGIS and our TopoToolbox script is in handling local minimum. ArcGIS uses the filling method, which ~~completely~~ fills depressions. In TopoToolbox, we used the carving method, which carves a channel through local depressions to allow the river to flow out. Therefore, although both used the same DEM, there are still slight differences in the positions of drainage system (Fig. 8a-d). Except for a 1.6% difference in river length, all the remaining seven parameters exhibit ~~variances of less than 0.8%~~ less than 0.8% variances. Basin90m's river channels and drainage divides are more ~~elaborate~~ detailed than those extracted by ArcGIS (Fig. 8b-d), resulting in longer ~~river~~ channels. Therefore, in Moche basin, Basin90m has higher spatial accuracy (Fig. 8b-d) and provides ~~more~~ accurate values for the morphometric metrics (Fig. 8e).

4.3 Accuracy of basin area

Drainage area is an important metric to evaluate the accuracy of drainage basin delineation. We compared drainage areas ~~from~~ Basin90m with those in HydroATLAS (Linke et al., 2019). We selected ten basins across the USA for comparison (Fig. 9a). These drainage basins are evenly distributed in the east-west direction across the North American continent, thus spanning diverse geological regimes, terrains, climates, and vegetation environments. Basins #1-5 are located in the west with arid climate, steep topography, and relatively active tectonics. Basins #6-10 are located in the east with low topographic relief but more precipitation and vegetation.

Due to the difference in DEM resolution (90 m for Basin90m and 500 m for HydroATLAS) and algorithms for delineating basin boundaries, the drainage divides from the two databases do not always coincide (Fig. 9b). But the two databases are consistent without substantial discrepancies (Fig. 9c). We quantified their difference using absolute relative error. The absolute relative error is the absolute value of the area difference between Basin90m and HydroATLAS as a ratio to HydroATLAS. Smaller values indicate higher agreement between the two datasets. Except for a 1.2% absolute relative error for basin #4, the values for the other nine basins are below 0.8%. The average absolute relative error for the ten basins is 0.47% (Fig. 9c). This slight discrepancy is acceptable given the nearly five-fold difference in DEM resolution and variations in algorithms for basin delineation.

In summary, the validation results based on ~~the~~ Peruvian Moche River basin and ten representative basins ~~spanning east-west~~ across the USA indicate that Basin90m has a high resolution for basin boundaries and river channels. Besides, the morphological parameters exhibit high accuracy.

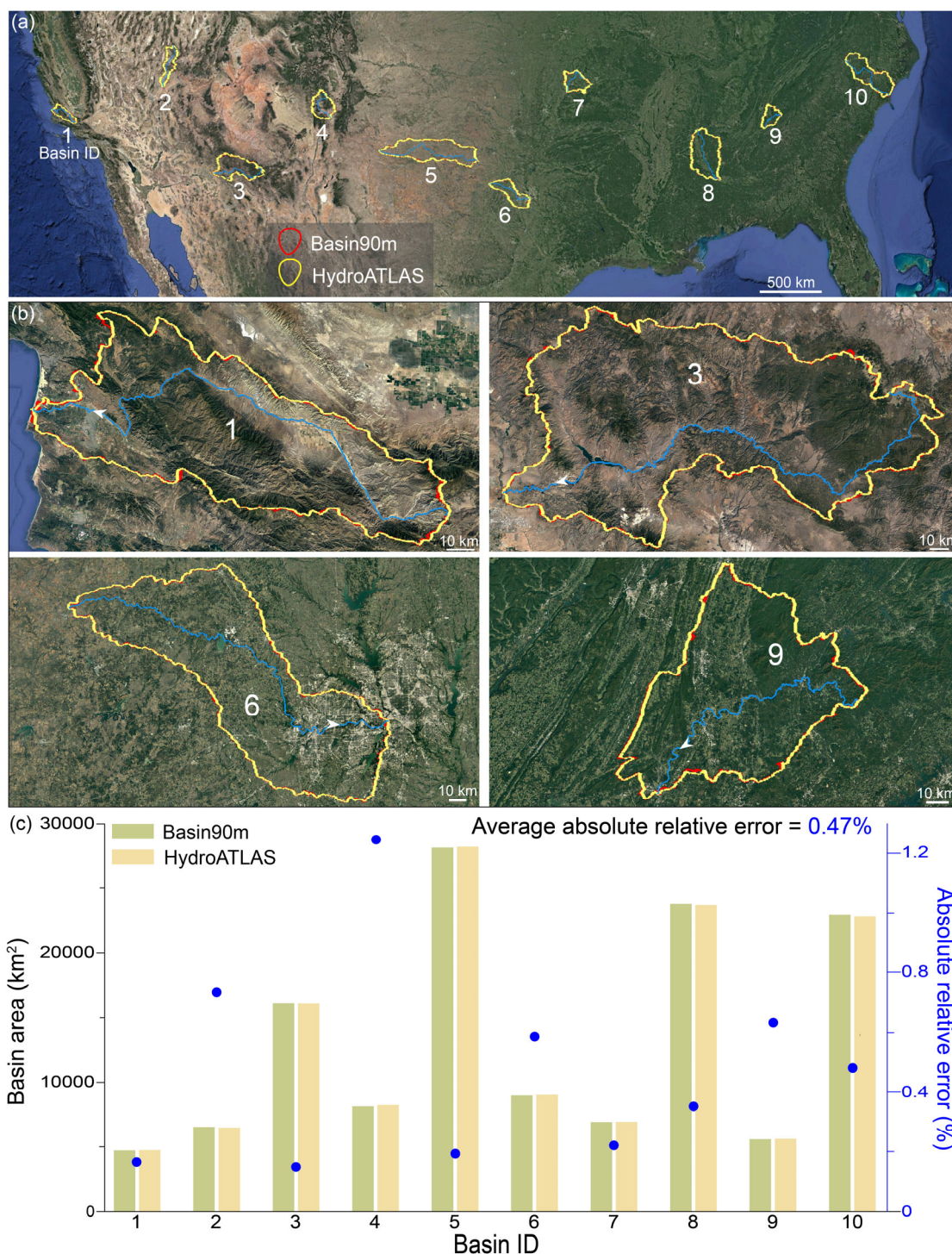


Figure 9. The accuracy of basin area in Basin90m, compared against HydroATLAS (Linke et al., 2019). The stream order for all the ten representative drainage basins is four. (a) Google Earth image shows ten example drainage basins in the USA. (b) Enlarged images of four representative basins. Basin #1 has an ocean outlet. Basin #3 is situated in a high-altitude arid region. Basin #6 features a flat terrain and encompasses a major city (Dallas). Basin #9 is located in a tectonically active folded region. (c) Comparison of basin areas between Basin90m and HydroATLAS.

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4.3—4 Limitations

555 Although Basin90m has improved ~~the~~ spatial accuracy of drainage systems compared to HydroSHEDS, it still has some limitations. First, the basins and rivers in Basin90m are entirely based on DEM analysis. Therefore, the measurement and processing errors of the DEM itself can affect the spatial accuracy of Basin90m. Second, in nature, rivers can flow in any direction. ~~But~~ However, the D8 method was used to calculate flow direction, which reduces computational
560 complexity but limits the flow to only eight directions. Third, after obtaining 667629 global basins and their longest rivers, no manual removal of false rivers has been conducted, especially in flat areas. Users must consider these limitations when using and interpreting the results of Basin90m, particularly in flat ~~regions and in~~ areas with severe human modifications.

565 5 Data availability

Basin90m is freely available at <https://dataservices.gfz-potsdam.de/panmetaworks/review/218da89b6b9388ce9161dc8462bd0ad65904300f2ec0fecfb53b45788c5614aa/> (He et al., 2023). Basin90m contains 667629 drainage basins with areas ~~larger than~~ over 50 km². Each basin is accompanied by its longest river channel, from drainage divide to river mouth. Basins and rivers
570 are stored in ESRI shapefile format, which can be opened and edited using GIS software (e.g., QGIS and ArcGIS) and Python libraries (e.g., GeoPandas).

 The data is grouped by six continents. Each continent contains multiple stream orders. For example, the ~~basin files for Europe~~ European basin files consist of eight shapefiles corresponding to stream orders 1-8. The filenames of basins and rivers include the stream order and continent information. For instance, “Africa_Basin_5.shp” contains all African basins with
575 a stream order of 5. The eight parameters (~~Fig. 2e~~) describing the size and shape of drainage systems (Fig. 2c) were stored in the attribute tables of ~~the~~ basin shapefiles. The shapefiles for global basins and rivers are 7.8 and 2.5 GB, respectively. The attribute tables of all global basins were merged into a single Excel file (Basin90m.xlsx).

580 6 Code availability

585 The Matlab script (Basin90m.m) used to generate Basin90m and a user guide are freely available at <https://dataservices.gfz-potsdam.de/panmetaworks/review/218da89b6b9388ce9161dc8462bd0ad65904300f2ec0fecfb53b45788c5614aa/>. Note that before running this code, one

needs to install TopoToolbox (Schwanghart and Scherler, 2014). The link <https://topotoolbox.wordpress.com/topotoolbox/> provides installation and usage instructions for TopoToolbox.

7 Conclusions

~~Here w~~We present Basin90m, a global dataset of the shape of drainage systems. Utilizing a 90-m resolution DEM, we extracted 667629 drainage basins with an area ~~larger than~~over 50 km². Each basin contains ~~only one~~ longest river channel, extending from the upstream drainage divide to the downstream river mouth. Basin90m provides information on the size, shape, hierarchy, and topography of drainage basins, as well as the length and sinuosity of river channels. Compared to the published datasets, Basin90m offers a higher resolution; ~~it~~ includes the shape of basins and rivers; ~~it~~ and excludes drainage systems with over half of their area located in lakes ~~and or~~ sandy deserts.

~~First, w~~We presented the variations among different stream orders regarding quantity, size, and shape of drainage systems. The number of basins decreased from 521857 for the first order to 3 for the ninth order. In contrast, increasing stream order ~~increased~~raised the average basin area, river length, and sinuosity. ~~Second, w~~We displayed the probability and spatial distribution of the eight parameters. The most notable feature is that numerous narrow and steep basins are distributed in the margins of ~~the~~ Tibetan Plateau. We then demonstrated the correlations among the eight parameters. The highest correlation coefficient of 0.97 was found between basin length and river length, indicating the coevolution of rivers and basins in nature. Using the basin area and river length from Basin90m, we fitted a global Hack's law as $L=2.1A^{0.54}$.

To validate the accuracy of Basin90m, we compared it with multiple data sources using the Moche ~~Basin~~ basin in Peru ~~and ten drainage basins across the North American continent as as an~~ examples. ~~This~~ The data compared with Basin90m include: basin boundaries, river locations and lengths provided by HydroSHEDS; basin area, slope, and elevation from HydroATLAS; basin and river locations from Google Earth images; basin boundaries and area, river locations and lengths, and drainage system shape metrics extracted using ArcGIS. ~~comparison included HydroSHEDS for river and basin boundaries, Google Earth images, drainage systems extracted using ArcGIS, and HydroATLAS for basin area, elevation, and slope.~~ The results showed that Basin90m exhibited the highest spatial accuracy. Furthermore, the difference ~~between of~~ morphological parameters ~~between in~~ Basin90m and other data sources is typically less than 1%. These validations confirm the ~~reliability and~~ accuracy of Basin90m as a valuable resource for ~~studying and~~ analyzing drainage systems on a ~~local or~~ global scale.

625 Author contributions

CH, CJY, and JMT conceived the initial idea. CH performed calculations and data processing, wrote a first draft, and made the figures. CJY built the code to extract Basin90m from DEM, with inputs from RFO. JMT designed method for measuring the length and width of drainage basin. RFO collected the global DEM. GSdQ proposed using the aridity index to ~~automatically~~
630 ~~classify sandy deserts~~classify sandy deserts. JB, HT, SG, and XPY participated in data processing and analysis. All authors contributed to the conceptualization, discussion, data collection, and editing of all components of Basin90m and this manuscript.

Competing interests

635 The contact author has declared that none of the authors has any competing interests.

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References

- Allen, G. H. and Pavelsky, T. M.: Global extent of rivers and streams, *Science*, 361, 585–588, <https://doi.org/10.1126/science.aat0636>, 2018.
- 650 Amatulli, G., Marquez, J. G., Sethi, T., Kiesel, J., Grigoropoulou, A., Üblacker, M. M., Shen, L. Q., and Domisch, S.: Hydrography90m: a new high-resolution global hydrographic dataset, *Earth Syst. Sci. Data*, 14, 4525–4550, <https://doi.org/10.5194/essd-14-4525-2022>, 2022.
- Bennett, S. J. and Liu, R.: Basin self-similarity, Hack's law, and the evolution of experimental rill networks, *Geology*, 44, 35–38, <https://doi.org/10.1130/G37214.1>, 2016.
- 655 Biron, P. M., Buffin-Belanger, T., Larocque, M., Chone, G., Cloutier, C. A., Ouellet, M. A., Demers, S., Olsen, T., Desjarlais, C., and Eyquem, J.: Freedom space for rivers: a sustainable management approach to enhance river resilience, *Environ. Manage.*, 54, 1056–1073, <https://doi.org/10.1007/s00267-014-0366-z>, 2014.
- Castelltort, S., Goren, L., Willett, S. D., Champagnac, J.-D., Herman, F., and Braun, J.: River drainage patterns in the New Zealand Alps primarily controlled by plate tectonic strain, *Nat. Geosci.*, 5, 744–748, <https://doi.org/10.1038/ngeo1582>, 2012.
- 660 [Cook, K. L., Rekapalli, R., Dietze, M., Pilz, M., Cesca, S., Rao, N. P., Srinagesh, D., Paul, H., Metz, M., Mandal, P., Suresh, G., Cotton, F., Tiwari, V. M., and Hovius, N.: Detection and potential early warning of catastrophic flow events with regional seismic networks. *Science*, 374, 87–92, <https://doi.org/10.1126/science.abj1227>, 2021.](https://doi.org/10.1126/science.abj1227)
- 665 [Datry, T., Boulton, A. J., Fritz, K., Stubbington, R., Cid, N., Crabot, J. and Tockner, K.: Non-perennial segments in river networks. *Nat. Rev. Earth Environ.*, 4, 815–830, <https://doi.org/10.1038/natrev-earth-environ.4.815>.](https://doi.org/10.1038/natrev-earth-environ.4.815)

- <https://doi.org/10.1038/s43017-023-00495-w>, 2023.
- 670 Farr, T. G., Rosen, P. A., Caro, E., Crippen, R., Duren, R., Hensley, S., Kobrick, M., Paller, M., Rodriguez, E., Roth, L., Seal, D., Shaffer, S., Shimada, J., Umland, J., Werner, M., Oskin, M., Burbank, D., and Alsdorf, D.: The Shuttle Radar Topography Mission, *Rev. Geophys.*, 45, RG2004, <https://doi.org/10.1029/2005RG000183>, 2007.
- Gamo, M., Shinoda, M., and Maeda, T.: Classification of arid lands, including soil degradation and irrigated areas, based on vegetation and aridity indices, *Int. J. Remote Sens.*, 34, 6701-6722, <https://doi.org/10.1080/01431161.2013.805281>, 2013.
- 675 Guth, P. L.: Drainage basin morphometry: a global snapshot from the shuttle radar topography mission, *Hydrol. Earth Syst. Sci.*, 15, 2091-2099, <https://doi.org/10.5194/hess-15-2091-2011>, 2011.
- Habousha, K., Goren, L., Nativ, R., and Gruber, C.: Plan-form evolution of drainage basins in response to tectonic changes: Insights from experimental and numerical landscapes, *J. Geophys. Res.: Earth Surf.*, 128, e2022JF006876, <https://doi.org/10.1029/2022JF006876>, 2023.
- 680 Hack, J. T.: Studies of longitudinal stream profiles in Virginia and Maryland, United States Geological Survey, <https://pubs.usgs.gov/pp/0294b/report.pdf>, 1957.
- He, C., Yang, C. J., Turowski, J. M., Rao, G., Roda-Boluda, D. C., and Yuan, X. P.: Constraining tectonic uplift and advection from the main drainage divide of a mountain belt, *Nat. Commun.*, 12, 544, <https://doi.org/10.1038/s41467-020-20748-2>, 2021a.
- 685 He, C., Yang, C. J., Rao, G., Roda-Boluda, D. C., Yuan, X. P., Yang, R., Gao, L., and Zhang, L.: Landscape response to normal fault linkage: Insights from numerical modeling, *Geomorphology*, 388, 107796, <https://doi.org/10.1016/j.geomorph.2021.107796>, 2021b.
- He, C., Yang, C. J., Turowski, J. M., Ott, R. F., Braun, J., Tang, H., Ghantous, S., Yuan, X. P., Stucky de Quay, G.: Basin90m, a new global drainage basin dataset. GFZ Data Services, <https://doi.org/10.5880/GFZ.4.6.2023.004>, 2023.
- 690 Hou, J., Van Dijk, A. I. J. M., Renzullo, L. J., and Larraondo, P. R.: GloLakes: a database of global lake water storage dynamics from 1984 to present derived using laser and radar altimetry and optical remote sensing, *Earth Syst. Sci. Data*, <https://doi.org/10.5194/essd-2022-266>, 2022.
- 695 Ielpi, A., Lapôtre, M. G. A., Finotello, A., and Roy-Léveillé, P.: Large sinuous rivers are slowing down in a warming Arctic, *Nat. Clim. Change*, 13, 375-381, <https://doi.org/10.1038/s41558-023-01620-9>, 2023.
- [Kirchner, J. W., Feng, X., and Neal, C.: Catchment-scale advection and dispersion as a mechanism for fractal scaling in stream tracer concentrations, *J. Hydrol.*, 254, 82-101, \[https://doi.org/10.1016/S0022-1694\\(01\\)00487-5\]\(https://doi.org/10.1016/S0022-1694\(01\)00487-5\), 2001.](https://doi.org/10.1016/S0022-1694(01)00487-5)
- 700 Lehner, B. and Grill, G.: Global river hydrography and network routing: baseline data and new approaches to study the world's large river systems, *Hydrol. Processes*, 27, 2171-2186, <https://doi.org/10.1002/hyp.9740>, 2013.
- Lehner, B., Verdin, K., and Jarvis, A.: New global hydrography derived from spaceborne elevation data, *Eos, Trans., Am. Geophys. Union*, 89, 93-104, <https://doi.org/10.1029/2008EO100001>, 2008.
- 705 Lin, P., Pan, M., Wood, E. F., Yamazaki, D., and Allen, G. H.: A new vector-based global river network dataset accounting for variable drainage density, *Sci. Data*, 8, 28, <https://doi.org/10.1038/s41597-021-00819-9>, 2021.
- Lindsay, J. B.: Efficient hybrid breaching-filling sink removal methods for flow path enforcement in digital elevation models, *Hydrol. Processes*, 30, 846-857, <https://doi.org/10.1002/hyp.10648>, 2016.
- 710 Linke, S., Lehner, B., Ouellet Dallaire, C., Ariwi, J., Grill, G., Anand, M., Beames, P., Burchard-Levine, V., Maxwell, S., Moidu, H., Tan, F., and Thieme, M.: Global hydro-environmental sub-basin and river reach characteristics at high spatial resolution, *Sci. Data*, 6, 283, <https://doi.org/10.1038/s41597-019-0300-6>, 2019.
- Luo, W., Howard, A. D., Craddock, R. A., Oliveira, E. A., and Pires, R. S.: Global spatial distribution of Hack's Law exponent on Mars consistent with early arid climate, *Geophys. Res. Lett.*, 50, e2022GL102604, <https://doi.org/10.1029/2022GL102604>, 2023.
- 715 Mantilla, R., Troutman, B. M., and Gupta, V. K.: Testing statistical self-similarity in the topology of river networks, *J. Geophys. Res.*, 115, F03038, <https://doi.org/10.1029/2009JF001609>, 2010.
- Masutomi, Y., Inui, Y., Takahashi, K., and Matsuoka, Y.: Development of highly accurate global polygonal drainage basin data, *Hydrol. Processes*, 23, 572-584, <https://doi.org/10.1002/hyp.7186>, 2009.
- 720 [Matthews, W. J., and Robison, H. W.: Influence of drainage connectivity, drainage area and regional species richness on fishes of the interior highlands in Arkansas. *Am. Midl. Nat.*, 139, 1-19, \[https://doi.org/10.1674/0003-0031\\(1998\\)139\\[0001:IODCDA\\]2.0.CO;2\]\(https://doi.org/10.1674/0003-0031\(1998\)139\[0001:IODCDA\]2.0.CO;2\), 1998.](https://doi.org/10.1674/0003-0031(1998)139[0001:IODCDA]2.0.CO;2)
- [Maurer, J. M., Schaefer, J. M., Russell, J. B., Rupper, S., Wangdi, N., Putnam, A. E., and Young, N.: Seismic observations, numerical modeling, and geomorphic analysis of a glacier lake outburst flood in the Himalayas. *Sci. Adv.*, 6, eaba3645, <https://doi.org/10.1126/sciadv.aba3645>, 2020.](https://doi.org/10.1126/sciadv.aba3645)
- 725 McEwan, E., Stahl, T., Howell, A., Langridge, R., and Wilson, M.: Coseismic river avulsion on surface rupturing faults: Assessing earthquake-induced flood hazard, *Science*, 9, eadd2932, <https://doi.org/10.1126/sciadv.add293>, 2023.
- 730 Messenger, M. L., Lehner, B., Grill, G., Nedeva, I., and Schmitt, O.: Estimating the volume and age of water

- stored in global lakes using a geo-statistical approach, *Nat. Commun.*, 7, 13603, <https://doi.org/10.1038/ncomms13603>, 2016.
- Montgomery, D. R. and Dietrich, W. E.: Channel initiation and the problem of landscape scale, *Science*, 255, 826-830, <https://doi.org/10.1126/science.255.5046.826>, 1992.
- 735 Mueller, J. E.: Re-evaluation of the relationship of master streams and drainage basins, *Geol. Soc. Am. Bull.*, 83, 3471-3474, [https://doi.org/10.1130/0016-7606\(1972\)83\[3471:ROTR0M\]2.0.CO;2](https://doi.org/10.1130/0016-7606(1972)83[3471:ROTR0M]2.0.CO;2), 1972.
- Nagayama, S. and Nakamura, F.: The significance of meandering channel to habitat diversity and fish assemblage: a case study in the Shibetsu River, northern Japan, *Limnology*, 19, 7-20, <https://doi.org/10.1007/s10201-017-0512-4>, 2017.
- 740 O'Malley, C. P. B.: Quantitative analysis of river profiles and fluvial landscapes, Doctoral dissertation, University of Cambridge, <https://doi.org/10.17863/CAM.51663>, 2020.
- Palmer, P. I., Wainwright, C. M., Dong, B., Maidment, R. I., Wheeler, K. G., Gedney, N., Hickman, J. E., Madani, N., Folwell, S. S., Abdo, G., Allan, R. P., Black, E. C. L., Feng, L., Gudoshava, M., Haines, K., Huntingford, C., Kilavi, M., Lunt, M. F., Shaaban, A., and Turner, A. G.: Drivers and impacts of Eastern African rainfall variability, *Nat. Rev. Earth Environ.*, 4, 254-270, <https://doi.org/10.1038/s43017-023-00397-x>, 2023.
- 745 Penido, J. C., Fassett, C. I., and Som, S. M.: Scaling relationships and concavity of small valley networks on Mars, *Planet. Space Sci.*, 75, 105-116, <https://doi.org/10.1016/j.pss.2012.09.009>, 2013.
- Rhoads, B. L., Schwartz, J. S., and Porter, S.: Stream geomorphology, bank vegetation, and three-dimensional habitat hydraulics for fish in midwestern agricultural streams, *Water Resour. Res.*, 39, 1218, <https://doi.org/10.1029/2003WR002294>, 2003.
- Sassolas-Serrayet, T., Cattin, R., and Ferry, M.: The shape of watersheds, *Nat. Commun.*, 9, 3791, <https://doi.org/10.1038/s41467-018-06210-4>, 2018.
- 755 Schwanghart, W., Groom, G., Kuhn, N. J., and Heckrath, G.: Flow network derivation from a high resolution DEM in a low relief, agrarian landscape, *Earth Surf. Process. Landf.*, 38, 1576-1586, <https://doi.org/10.1002/esp.3452>, 2013.
- Schwanghart, W. and Scherler, D.: Bumps in river profiles: uncertainty assessment and smoothing using quantile regression techniques, *Earth Surf. Dyn.*, 5, 821-839, <https://doi.org/10.5194/esurf-5-821-2017>, 2017.
- 760 Schwanghart, W. and Scherler, D.: Short communication: TopoToolbox 2 – MATLAB-based software for topographic analysis and modeling in Earth surface sciences, *Earth Surf. Dyn.*, 2, 1-7, <https://doi.org/10.5194/esurf-2-1-2014>, 2014.
- [Shelef, E.: Channel profile and plan-view controls on the aspect ratio of river basins, *Geophys. Res. Lett.*, 45, 11712-11721, <https://doi.org/10.1029/2018GL080172>, 2018.](https://doi.org/10.1029/2018GL080172)
- 765 Shen, X., Anagnostou, E. N., Mei, Y., and Hong, Y.: A global distributed basin morphometric dataset, *Sci. Data*, 4, 160124, <https://doi.org/10.1038/sdata.2016.124>, 2017.
- Sikder, M. S., Wang, J., Allen, G. H., Sheng, Y., Yamazaki, D., Song, C., Ding, M., Crétaux, J.-F., and Pavelsky, T. M.: Lake-TopoCat: A global lake drainage topology and catchment database, *Earth Syst. Sci. Data*, 15, 3483-3511, <https://doi.org/10.5194/essd-15-3483-2023>, 2023.
- 770 Som, S. M., Montgomery, D. R., and Greenberg, H. M.: Scaling relations for large Martian valleys, *J. Geophys. Res.*, 114, E02005, <https://doi.org/10.1029/2008JE003132>, 2009.
- Sreedevi, P. D., Owais, S., Khan, H. H., and Ahmed, S.: Morphometric analysis of a watershed of South India using SRTM data and GIS, *J. Geol. Soc. India*, 73, 543-552, <https://doi.org/10.1007/s12594-009-0038-4>, 2009.
- 775 Strahler, A. N.: Quantitative analysis of watershed geomorphology, *Eos, Trans., Am. Geophys. Union*, 38, 913-920, <https://doi.org/10.1029/TR038i006p00913>, 1957.
- Strong, C. M. and Mudd, S. M.: Explaining the climate sensitivity of junction geometry in global river networks, *PNAS*, 119, e2211942119, <https://doi.org/10.1073/pnas.2211942119>, 2022.
- 780 Tarboton, D. G.: A new method for the determination of flow directions and upslope areas in grid digital elevation models, *Water Resour. Res.*, 33, 309-319, <https://doi.org/10.1029/96WR03137>, 1997.
- Toyne, J. M., White, C. D., Verano, J. W., Uceda Castillo, S., Millaire, J. F., and Longstaffe, F. J.: Residential histories of elites and sacrificial victims at Huacas de Moche, Peru, as reconstructed from oxygen isotopes, *J. Archaeol. Sci.*, 42, 15-28, <https://doi.org/10.1016/j.jas.2013.10.036>, 2014.
- 785 Tu, T., Comte, L., and Ruhi, A.: The color of environmental noise in river networks, *Nat. Commun.*, 14, 1728, <https://doi.org/10.1038/s41467-023-37062-2>, 2023.
- USGS: HYDRO1k elevation derivative database, U.S. Geol. Surv., <https://doi.org/10.5066/F77P8WN0>, 2000.
- Verdin, K. L.: Hydrologic derivatives for modeling and applications (HDMA) database-A new global high-resolution database, U.S. Geol. Surv., Virginia, <https://doi.org/10.3133/ds1053>, 2017.
- 790 Vörösmarty, C. J., Fekete, B. M., Meybeck, M., and Lammers, R. B.: Global system of rivers: Its role in organizing continental land mass and defining land-to-ocean linkages, *Global Biogeochem. Cycles*, 14, 599-621, <https://doi.org/10.1029/1999GB900092>, 2000.
- [Yamazaki, D., Ikeshima, D., Sosa, J., Bates, P. D., Allen, G. H., and Pavelsky, T. M.: MERIT Hydro: A high-resolution global hydrography map based on latest topography dataset, *Water Resour. Res.*, 55, 5053-](https://doi.org/10.1029/2018WR023000)

~~5073, <https://doi.org/10.1029/2019WR024873>, 2019.~~

- 795 Yan, D., Li, C., Zhang, X., Wang, J., Feng, J., Dong, B., Fan, J., Wang, K., Zhang, C., Wang, H., Zhang, J.,
and Qin, T.: A data set of global river networks and corresponding water resources zones divisions v2,
Sci. Data, 9, 770, <https://doi.org/10.1038/s41597-019-0243-y>, 2022.
- Yi, R. S., Arredondo, Á., Stansifer, E., Seybold, H., and Rothman, D. H.: Shapes of river networks, Proc. R.
Soc. A 474, 20180081, <https://doi.org/10.1098/rspa.2018.0081>, 2018.
- 800 Yu, Z., Zhang, J., Wang, H., Zhao, J., Dong, Z., Peng, W., and Zhao, X.: Quantitative analysis of ecological
suitability and stability of meandering rivers, Front. Biosci. Landmark, 27, 42,
<https://doi.org/10.31083/j.fbl2702042>, 2022.
- Yuan, X. P., Jiao, R., Liu-Zeng, J., Dupont-Nivet, G., Wolf, S. G., Shen, X.: Downstream propagation of fluvial
erosion in Eastern Tibet. Earth Planet. Sci. Lett., 605, 118017, <https://doi.org/10.1016/j.epsl.2023.118017>,
805 2023.
- Zomer, R. J., Xu, J., and Trabucco, A.: Version 3 of the Global Aridity Index and Potential Evapotranspiration
Database, Sci. Data, 9, 409, <https://doi.org/10.1038/s41597-022-01493-1>, 2022.