GTWS-MLrec: Global terrestrial water storage reconstruction by machine learning from 1940 to present

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Abstract. Terrestrial water storage (TWS) includes all forms of water stored on and below the land surface, and is a key determinant of global water and energy budgets. However, TWS data from measurements by the Gravity Recovery and Climate Experiment (GRACE) satellite mission are only available from 2002, limiting global and regional understanding of the long-term trends and variabilities in the terrestrial water cycle under climate change. This study presents long-term (i.e., 1940-2022) and high-resolution (i.e., 0.25°) monthly time series of TWS anomalies over the global land surface. The reconstruction is achieved by using a set of machine learning models with a large number of predictors, including climatic and hydrological variables, land use/land cover data, and vegetation indicators (e.g., leaf area index). The outcome, machine learning-reconstructed TWS estimates (i.e., GTWS-MLrec), fits well with the GRACE/GRACE-FO measurements, showing high correlation coefficients and low biases in the GRACE era. We also evaluate GTWS-MLrec with other independent products such as the land-ocean mass budget, atmospheric and terrestrial water budget in 341 large river basins, and streamflow measurements at 10,168 gauges. The results show that our proposed GTWS-MLrec performs overall as well as or is more reliable than previous TWS datasets. Moreover, our reconstructions successfully reproduce the consequences of climate variability, such as strong El Niño events. GTWS-MLrec dataset consists of three reconstructions based on JPL, CSR and GSFC mascons, three detrended and de-seasonalized reconstructions, and six global average TWS series over land areas, both with and without Greenland and Antarctica. Along with its extensive attributes, GTWS_MLrec can support a wide range of geoscience applications such as better understanding the global water budget, constraining and evaluating hydrological models, climate-carbon coupling, and water resources management. GTWS-MLrec is available on Zenodo through https://zenodo.org/record/8187432 (Yin et al., 2023c).

1 Introduction

Information on global water cycle dynamics is crucial for the monitoring of water-related programs as well as scientific investigations such as understanding the spatiotemporal variability in terrestrial freshwater availability (Lettenmaier and
40 Famiglietti, 2006). Terrestrial Water Storage (TWS) includes all components of water reservoirs (i.e., ice, snow, wetlands, lakes, rivers, soil moisture and groundwater) on and below the continental land surface, and is a necessary element to close the terrestrial water budget, which enables to balance evapotranspiration, precipitation, and runoff at both regional and global scales (Pokhrel et al., 2012; Kusche et al., 2016). As an essential driver of global water and energy budgets, TWS is also highly sensitive to global climate change, and therefore has been widely employed to assess the impacts of large-scale
45 hydrological extremes (e.g., droughts and floods) on socioeconomic systems and ecosystem sustainability across the warming planet (Yin et al., 2023a).

TWS fluctuations can be altered by natural processes, anthropogenic climate warming and human activities (Pokhrel et al., 2012; Felfelani et al., 2017). For example, the El Niño–Southern Oscillation (ENSO), which represents natural variability
50 in ocean and atmospheric circulation, can alter the anomalies in atmospheric water vapor transport and thus leading to regional precipitation deficit or excess (Ni et al., 2017; Zhang et al., 2023). Anthropogenic climate warming has intensified the global water cycle, with heavier precipitation and diminishing snowmelt, resulting in significant spatial heterogeneity of TWS fluctuations globally (Yin et al., 2022a; Gu et al., 2023). Fluctuations in TWS have been widely reported to correlate with a broad range of natural phenomena, such as changes in global ocean mass, alternation of carbon uptake by terrestrial
55 ecosystems, and the movement of the rotational axis of Earth (Kim et al., 2019; Humphrey and Gudmundsson, 2019). In addition to the climate variability, TWS is also highly impacted by intense anthropogenic activities, such as rapid urbanization, irrigation, reservoir operation, groundwater depletion, water diversion projects, and the recent ice/glacier retreat (Pokhrel et al., 2012; Long et al., 2020; Jacob et al., 2021). Given the large number of factors that are correlated with TWS, numerous studies have been devoted to understanding TWS evolution and drivers at a catchment, regional or global
60 scale (e.g., Wang et al., 2020; Zhao et al., 2021).

In March 2002, the Gravity Recovery and Climate Experiment (GRACE) satellite was launched which started to provide a continuously direct device to monitor the global TWS variations with an unprecedented spatiotemporal resolution (Wahr et al., 2004). After one-year cease of GRACE's monitoring mission, the successor satellite (i.e., GRACE Follow-On, FO) was
65 launched in May 2018. Although the GRACE/GRACE-FO satellites have effectively measured global water cycle dynamics since 2002, there was very limited global understanding or monitoring of TWS beyond the GRACE era or during the gap period between GRACE and GRACE-FO. In recent years, long-term TWS data has become a growing requirement for a

wide range of climatic or hydrological fields, such as constraining the ocean mass budget, improving global hydrological

models, exploring climate signal fluctuations, and understanding hydrological extremes and their impacts on ecosystem

70   productivity (e.g., Chambers et al., 2016; Markonis et al., 2018; Pokhrel et al., 2021; Yin et al., 2023a).

Several statistical methods and hydrological models have been used to retrospectively reconstruct TWS beyond the GRACE

era at either a catchment or regional scale (Humphrey et al., 2017; Ahmed et al., 2019; Sun et al., 2020). Among these, very

few studies have been devoted to reconstructing TWS at a global scale, and the GRACE-REC product (0.5° resolution)

75   reconstructed by Humphrey and Gudmundsson (2019) has received the most attention in hydrological studies. GRACE-REC

exhibits an overall better performance than a set of global hydrological models, and well represents the variations in water

storage due to climate change during the last century. However, GRACE-REC does not include the seasonal TWS cycle,

and most of the reconstruction dataset do not extend prior to 1979. More recently, Li et al. (2020; 2021) developed a three-

stage approach to reconstruct a so-called GRACE-like TWS during 1979–2020 at a global scale, and found that their

80   reconstructions with 0.5° resolution agree reliably with the GRACE/GRACE-FO measurements and represent strong water

anomalies during El Niño years. While these approaches have provided important reference data linking TWS and multiple

predictors, the predictors have typically been restricted to climatic variables, and rarely consider land use/land cover data

and vegetation indicators such as Leaf Area Index (LAI). Further, to the best of our knowledge, none of the previous studies

has yet reconstructed TWS prior to 1979 at a spatial resolution finer than 0.5° at a global scale.

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The primary objective of this work is to fill the existing gap in historical TWS data and provide long-term (i.e., 1940-2022)

and high-resolution (i.e., 0.25°) monthly time series of TWS anomalies over global land areas. The reconstruction is achieved

by a set of machine learning models using broad input drivers, including climatic and hydrological variables, land use/land

cover data, and vegetation indicators. The machine learning models are trained by GRACE/GRACE-FO measurements, and

90   our reconstructions, named as GTWS-MLrec, agree well with the observations. We also evaluate GTWS-MLrec

reconstructions with numerous independent products/methods such as the land-ocean water mass budget, atmospheric and

terrestrial water budget over large catchments, and *in situ* river streamflow observations at 10,168 gauges. In addition, our

reconstructions accurately reproduce the effects of climate anomalies such as strong El Niño events. Overall, our proposed

GTWS-MLrec performs as well as, or better than previous TWS reconstruction datasets in most conditions.

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2 Data and methods

2.1 GRACE/GRACE-FO measurements

Three different GRACE/GRACE-FO solutions based on mass concentration technique are employed, which cover the monthly TWS series over the period of 2002-2022. This newly developed algorithm provides estimations of mass variations over small and predefined regions, which are briefly referred to as mascons (Watkins et al., 2015). The mascon solutions are usually better than the spherical harmonics-based products; for example, it reduced leakage due to increased signal amplitude, and it requires fewer or no postprocessing procedures (Scanlon et al., 2016). We use the latest mascon-based products from three international centres: Jet Propulsion Laboratory (JPL) of California Institute of Technology, the Centre for Space Research (CSR) at the University of Texas at Austin, and the Goddard Space Flight Centre (GSFC) of National Aeronautics and Space Administration (NASA). The TWS products based on three mascon solutions are divergent, because they are produced by different processing methods and different employed models are employed for correcting the effect of glacial isostatic adjustment. The GRACE/GRACE-FO TWS estimates are employed to train our machine learning model to obtain the long-term reconstruction. The three mascon solutions have different spatial resolution due to divergent processing methods (Table 1). Although the GRACE/GRACE-FO datasets have some limitations such as spatial and temporal coarseness, they are the best available training data even if their measurements can be further improved.

2.2 Inputs for the machine learning models

Numerous meteorological or hydrological variables were identified as important elements in TWS reconstructions by previous works (e.g., Sun et al., 2020; Li et al., 2021). Here, we use four types of predictors to feed the machine learning model, including (1) eleven meteorological elements from the fifth generation European Centre for Medium-Range Weather Forecasts re-analysis (ERA5), (2) two hydrological variables from ERA5, (3) land use/land cover data, and (4) vegetation indicators, i.e., LAI and Solar-induced fluorescence (SIF). The eleven meteorological variables consist of 2-meter temperature (℃), near-surface specific humidity (kg kg^{-1}) and relative humidity (%), snowfall (mm), vertically integrated moisture convergence (kg m^{-2} s^{-1}), precipitation (mm), 10-meter wind speed (m s^{-1}), surface downward short-wave and long-wave radiation (W m^{-2}), evaporation (mm), and cloud cover (%). These variables are selected due to their representation of water flux (e.g., precipitation and moisture convergence), energy flux (e.g., temperature, radiation and cloud cover) and other processes involving water-energy transport (i.e., wind speed). Moisture convergence is represented by the negative value of vertically integrated moisture divergence in ERA5, and the cloud cover data refers to the proportion of a grid covered by clouds occurring in the lower levels of the troposphere.

The near-surface relative humidity (RH) and specific humidity (SH) are not currently available in the ERA5 monthly

ensemble dataset, which can be estimated by using 2-meter temperature (T_{2m}), dewpoint temperature (T_{dew}) and air pressure (pr). The Clausius-Clapeyron relationship describes the dependence of atmospheric saturation vapor pressure on air temperature as follows:

$$e_{sat}(T) = e_{s0}\, exp\left[\frac{L_v}{R_v}(\frac{1}{T_0} - \frac{1}{T})\right] \tag{1}$$

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where T and e_{sat} indicate the near-surface air temperature (°C) and saturation vapor pressure (Pa), respectively; L_v and R_v refer to the latent heat of vaporization (2.5×10^6 J kg^{-1}) and vapor gas constant (461 J kg^{-1} K^{-1}), respectively; $T_0 = 273.15$ K and $e_{s0} = 611$ Pa are both integration constants.

135 The T_{dew} denotes the temperature above which the air moisture will be saturated under constant water vapor content and pressure; therefore, it can characterize the actual atmospheric water vapor availability. The RH can be calculated by substituting T_{2m} and T_{dew} into Eq. (1) as $RH = e_{sat}(T_{dew})/e_{sat}(T_{2m})$.

SH represents the mass contribution of water vapor to the total air mixture, which can be derived by pr and T_{dew}:

$$SH = 0.622\,\frac{e_{sat}(T_{dew})}{pr - 0.378 e_{sat}(T_{dew})} \tag{2}$$

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Hydrological variables such as runoff and soil moisture have been shown to be highly correlated with TWS (Sun et al., 2020; Yang et al., 2023), and therefore these two variables also served as predictors in the reconstruction model. For soil moisture, we use the average volumetric soil water of four layers weighted by the layer depth. Numerous studies have reported that land use/land cover change has substantial impacts on TWS, for example changes in impervious surface area

145 (ISA) due to urbanization play an essential role in driving TWS variability (Chen et al., 2018; Wang et al., 2020). To constrain the TWS by considering the effects of urbanization, we have extracted the ISA series from the latest FROM-GLC Plus product using Google Earth Engine. The pixel-based ISA is calculated from a multi-temporal (i.e., from daily to annual) and multi-resolution (i.e., ranging from sub-meter to 30 m) global land cover product (Yu et al. 2022). We use ISA to represent land cover changes to simplify the inputs of our machine learning model, which enables more efficient

150 inference beyond the training period. Vegetation and TWS also have a feedback relationship due to water-energy exchange through photosynthesis and respiration as well as vegetation's regulation of soil moisture (Yin et al., 2023a, b; Liu et al., 2023). Therefore, we also use LAI from ERA5 and the recent satellite-based machine-learning-generated SIF dataset by Zhang et al. (2018) to train our machine learning models. For detailed information about the inputs in machine learning model, please refer to Table 1.

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Table 1. Main data inputs for the machine learning models

Products		Data period	Spatial resolution	Reference
GRACE/GRACE-FO TWS	JPL-RL06M mascons	April 2002 to November 2022	0.5°	Watkins et al. (2015)
	GSFC RL06 v1.0 mascon solution	April 2002 - November 2022	0.5°	Loomis et al. (2019)
	CSR RL06 v02 mascon solution	April 2002 to December 2022	0.25°	Save et al. (2016)
Meteorological variables (EAR5)	2-meter temperature	January 1940 to December 2022	0.25°	Hersbach et al. (2020)
	2-meter dewpoint temperature	January 1940 to December 2022	0.25°	Hersbach et al. (2020)
	Surface air pressure	January 1940 to December 2022	0.25°	Hersbach et al. (2020)
	Snowfall	January 1940 to December 2022	0.25°	Hersbach et al. (2020)
	Vertically integrated moisture divergence	January 1940 to December 2022	0.25°	Hersbach et al. (2020)
	Precipitation	January 1940 to December 2022	0.25°	Hersbach et al. (2020)
	10-meter wind speed	January 1940 to December 2022	0.25°	Hersbach et al. (2020)
	Surface downward short-wave radiation	January 1940 to December 2022	0.25°	Hersbach et al. (2020)
	Surface downward long-wave radiation	January 1940 to December 2022	0.25°	Hersbach et al. (2020)
	Evaporation	January 1940 to December 2022	0.25°	Hersbach et al. (2020)
	Low cloud cover	January 1940 to December 2022	0.25°	Hersbach et al. (2020)
Hydrological variables (ERA5)	Volumetric soil water (0-289cm)	January 1940 to December 2022	0.25°	Hersbach et al. (2020)
	Runoff	January 1940 to December 2022	0.25°	Hersbach et al. (2020)
Land use/land cover data	Impervious surface area	January 1982 to December 2020	1 km	Yu et al. (2022)
Vegetation indicators	Leaf area index, high vegetation (ERA5)	January 1940 to December 2022	0.25°	Hersbach et al. (2020)
	Leaf area index, low vegetation (ERA5)	January 1940 to December 2022	0.25°	Hersbach et al. (2020)
	Solar-induced chlorophyll fluorescence	January 2000 to December 2022	0.05°	Zhang et al. (2018)

2.3 Independent evaluation datasets

Two most widely used global TWS reconstruction datasets (0.5° resolution) are used as a comparison, i.e., the GRACE-

160   REC dataset by Humphrey and Gudmundsson (2019) and the recent GRACE-like reconstructed TWS dataset (i.e., denoted

as GRL-REC dataset thereafter) by Li et al. (2022). The GRACE-REC dataset is produced based on two GRACE/GRACE-

FO solutions and three meteorological forcing products, and therefore releases six reconstructed datasets which includes 100

members within each data scheme. To facilitate the comparison, we average these six series to produce an ensemble average

reconstructed TWS product, as it has already been shown that the blended product captures the GRACE-measured TWS

165   dynamics well (Yin et al., 2023a). To evaluate the performance of our GTWS-MLrec reconstruction, we also try to close the

ocean mass budget by using the global mean sea level (GMSL) from altimeter data (Nerem et al., 2018) and steric height

estimates (GMSL_ster) based on measurements of Argo profiling floats (Levitus et al., 2012). In this study, we use total

steric level (i.e., the sum of thermosteric and halosteric sea level) of the 0-700 meter layer of the ocean. As the GMSL_ster dataset only provides seasonal series after 2005, we combine the seasonal GMSL_ster series and the running pentanal (i.e.,

170   5 years) series prior to 2005. We also use TWS variations from the Basin-scale water-balance dataset (BSWB) as independent reference dataset. The BSWB dataset derived monthly variations in large-scale TWS for 341 large catchments (i.e., area > 10 000 km²) worldwide by employing a hybrid atmospheric and terrestrial water budget technique (Hirschi and Seneviratne, 2017).

175   As the BSWB only provides TWS variation estimates during 1979-2015, we also try to evaluate the accuracy of TWS reconstructions by exploring their relationship with annual streamflow over a larger number of catchments. To achieve this goal, we gather daily river streamflow records from 1940 to the present at 22, 538 hydrological gauges from multiple sources. These records are sourced from a large combination of national and global data archives: (1) the Global Runoff Data Centre (GRDC); (2) the UK Centre for Ecology and Hydrology (UKCEH); (3) the Environment and Climate Change Canada

180   through the Water Survey of Canada (ECCC); (4) the U.S. Geological Survey National Water Information System (USGS); (5) the Australian Bureau of Meteorology (ABM); (6) the Brazilian National Water Agency (ANA); (7) the Ministry of Water Resources of China (MWRC); (8) the watershed management agencies affiliated with the MWRC; (9) the National Hydrological Information Centre of China; (10) Guangdong Provincial Bureau of Beijing River Administration, China; and (11) Chaohu Lake Research Institute, China. First, we omit those stations with changing measurement instrument or station

185   datum to keep data consistency. Second, we exclude gauges with less than 20 valid years of data (with >90% completeness for each year). Finally, we screen the catchments, excluding those with a catchment area larger than 10,000 km². Using the daily streamflow records, we calculate the monthly average runoff depth by considering streamflow observations and catchment area, where only those months with >90% daily data completeness are considered. For each year with 12 valid monthly values, we sum the values to derive a yearly runoff depth by considering catchment area. Overall, these filtering

190   steps leave 10,168 catchments with complete 10-year annual streamflow, which cover diverse climatic patterns and underlying surface conditions across the globe.

2.4 Machine learning-based TWS reconstruction method

The workflow of this reconstruction approach contains five modules, which is illustrated in Figure 1. First, we employ five

195   different machine learning models as candidates for TWS reconstruction. Second, eight input schemes are built for each machine learning model. Third, the random forest (RF) associated with a moving-windows approach is employed to determine the dominant variables in each input scheme for each machine learning model. Fourth, the best-performing machine learning model is selected in terms of the TWS simulation performance in the test period. Finally, the simulation

performance of the eight input schemes is rated, and the eight schemes are blended to produce the long-term TWS series at

200   each pixel.
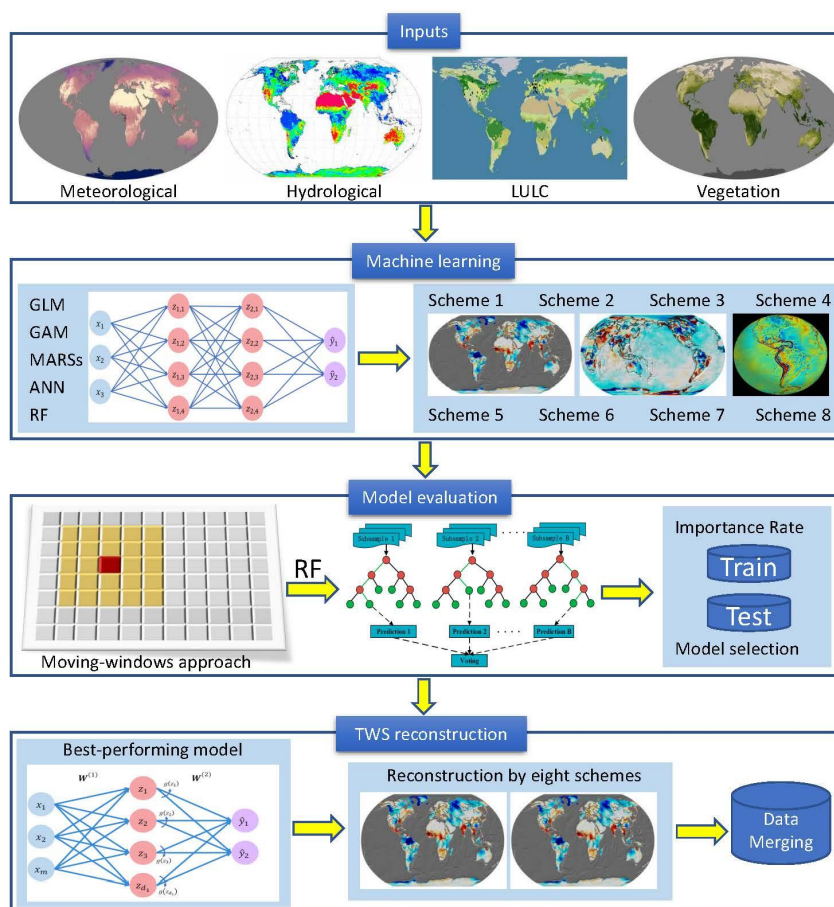
Figure 1. Flowchart of the machine learning-based TWS reconstruction approach.

The five data-driven (machine learning and statistical) models employed to establish the TWS reconstruction model are

205   listed as follows: multivariate adaptive regression splines (MARSs), Gaussian linear regression model (GLM), artificial

neural network (ANN), Gaussian generalised additive model (GAM), and a RF model. These five models employ the

regression-based techniques to characterize the relationship between the predictors (i.e., dependent climate/vegetation

variables) and the predictand (i.e., independent variables, TWS). The GLM is built within a parametric regression framework,

while the other four models are based on a non-parametric regression algorithm, where the functional relationship has the

210   feasibility of reliable adjustment to explore unusual or unexpected features (Shortridge et al., 2016; Singh et al., 2023). For

a more detailed illustration of these machine learning models, please refer to relevant references (e.g., Ghimire et al., 2021;

Herath et al., 2021; Li et al., 2021). For each machine learning model, eight different reconstruction schemes in terms of inputs are established: (1) scheme 1, containing all variables; (2) scheme 2, excluding LAI; (3) scheme 3, excluding SIF; (4) scheme 4, excluding LAI and SIF; (5) scheme 5, excluding LUCC; (6) scheme 6, excluding LUCC and LAI; (7) scheme 7,

215 excluding LUCC and SIF; (8) scheme 8, excluding LUCC, LAI and SIF. We establish eight different data schemes because some variables might be missing during some periods; by blending different data schemes, we will be able to achieve a more complete long-term series. In all eight construction schemes, a time lag of three months is considered for the inputs, i.e., the data at the current time step and in the previous 1-3 months are both employed to feed the machine learning model. For example, in Scheme 1, the input variables contain 64 time series of predictors as well as the GRACE/GRACE-FO TWS

220 observations.

As TWS is governed by divergent physical mechanisms under different underlying surface condition and climate patterns, the dominant variables for explaining the TWS may differ across different climatic regions (Yin et al., 2022a). Before establishing the TWS reconstruction model at each pixel, a moving-window nearest-neighbour approach is employed to

225 select the most important variables for each pixel and its immediate neighbours. To balance the size of data sample and model complexity, we use a moving-window size of as 5×5 for each pixel. The RF is employed to select the most important 60% of all candidate variables in each data scheme. The use of a moving window allows the model to be trained on a larger sample of data and to identify the most important candidate variables with greater consistency.

230 The eight data schemes for the five machine learning models are trained with GRACE/GRACE-FO data and multi-source inputs during 2002-2022 at each pixel, and their performance in simulating TWS is compared across the model-scheme combinations. Prior to building the machine learning models, all the data are normalised by using standard normalisation techniques to standardise the features on a common numerical scale. To evaluate the performance of the reconstruction models, a cross-validation method is employed and the entire dataset is randomly split into training and testing parts. The

235 training dataset (60 %) is employed to fit the models, while the remaining 40% of data is use to test the model accuracy (40 %). The R package "randomForest" is adopted to implement the RF-based analysis (Breiman, 2001).

In the reconstruction procedure, we first select the best-performing machine learning model based on the evaluation index (see Text S1) in scheme 8 during the test period, and then rate the simulation performance of the eight data schemes within

240 the best-simulating machine learning model. We select scheme 8 for selecting the best-simulating machine learning model, because this scheme contains the least volume of inputs and is more suitable for data extrapolation. After determining the best-performing model, we reconstruct the long-term TWS series by considering the scores of all the eight data schemes.

For example, the best-performing scheme of the selected machine learning model is used to produce the TWS reconstructions, and then the second-best performing scheme for the same model is used to fill any missing gaps of the former one. To

245   improve the capacity of the machine learning models to extrapolate TWS series beyond the calibration period, we employ all the observations to train our model in the reconstruction process. However, we also evaluate the performance of extrapolation by splitting train and test periods in Section 3.3. By blending the different data schemes with consideration of their simulation performance, the TWS series for the long-term period (i.e., 1940-2022) is fully reconstructed. Using three different training GRACE/GRACE-FO datasets (i.e., JPL, CSR and GSFC), we produce three different GTWS-MLrec

250   datasets. As many studies focus on exploring climate driven-TWS variability, we also produce three detrended and de-seasonalized TWS reconstruction series. To achieve this goal, we first employ the Seasonal-Trend decomposition method based on Loess (Rojo et al., 2017) to partition the GRACE-measured series into linear trends, seasonal and residuals. The residual component from the GRACE-measured temporal series is then reconstructed separately by using their empirical relationships as related to the potential predictors (i.e., components after detrending and de-seasoning from the input

255   variables). Therefore, we produce three full-component TWS datasets as well as three detrended and de-seasonalized TWS datasets across all global land areas for the period 1940-2022.

3 Data description and machine learning evaluation

3.1 Six GTWS-MLrec datasets

260   The GTWS-MLrec provides monthly TWS anomalies in units of millimeters of water (mm) during 1940-2022, with a spatial resolution of 0.25° across global land areas (including Greenland and Antarctica). Using three different training GRACE/GRACE-FO mascon solutions (Table 1), we produce three different GTWS-MLrec datasets. As some applications mainly focus on climate-driven TWS changes, we also provide de-seasonalized and detrended TWS anomalies, which are independently reconstructed by using the de-seasonalized and detrended GRACE/GRACE-FO dataset and inputs. It is

265   informative to note that these de-seasonalized and detrended TWS reconstructions are not necessarily systematically consistent with the reconstructed TWS datasets after the de-seasonalizing and detrending processes, because they are reconstructed by using independent machine learning models. Therefore, we provide a total of six reconstructed GTWS-MLrec datasets trained by three GRACE/GRACE-FO solutions.

270   ### 3.2 Global land mean TWS datasets

To provide data reference for a global-scale application, we also estimate a global average time series of the GTWS-MLrec TWS anomalies. The global mean TWS series is estimated by considering the area weight (i.e., the area of grid in different latitudes is considered), which includes all land areas with or without considering the Greenland and Antarctica (i.e., two

schemes are available). These global average TWS datasets can provide rich information as references in constraining global-

275   scale water dynamic, which is particularly suited for quantifying global land-ocean water budget. The global average TWS

has a unit of millimeters of water (mm). To convert millimeters to gigatons of water for mass budget applications, total land

areas of 148 940 000 km^2 and 132 773 914 km^2 can be used for each scheme, respectively.

3.3 Performance of machine learning model in extrapolating TWS beyond the calibration period

Before using the machine learning models for the TWS reconstruction, we evaluate their extrapolation performance by

280   randomly splitting the datasets into train and test periods during GRACE/GRACE-FO era. Eight metrics are used (Text S1),

including Pearson's Correlation Coefficient (PCC), Nash-Sutcliffe efficiency coefficient (NSE); Kling-Gupta Efficiency

coefficient (KGE), Coefficient of Determination (R^2), Root Mean square error (RMSE, unit: mm), normalized Root Mean

Square Error (nRMSE), Mean Absolute Percentage Error (MAPE), and the Percent bias (Pbias, unit: %). Figure 2 presents

the performance of the JPL-based reconstruction in data scheme 8 during the test period by comparison with the

285   GRACE/GRACE-FO measurements. The PCC, NSE, KGE and the R^2 approach 0.9 in most areas of the globe for the test

period, and the average nRMSE over the global land areas is about 0.1. The MAPE, RMSE and Pbias also suggest a good

performance. Comparing the five machine learning models, we find that the non-linear models (i.e., RF and ANN) show

superior performance than the linear models (i.e., GLM, GAM and MARSs). The GAM, MARSs and ANN also show a

relatively reliable performance in extrapolating TWS anomalies, with the global average PCC of about 0.7, and the GLM

290   shows the worst performance in most areas of the globe (Figure S1). We also compare the extrapolation performance for

GSFC and CSR solutions, and find similar conclusions, namely that the RF and ANN show better capacity in simulating

TWS anomalies in most areas of the globe. Overall, we are confident that the best-simulating machine learning model allows

to reliably extrapolate TWS series beyond the calibration period.

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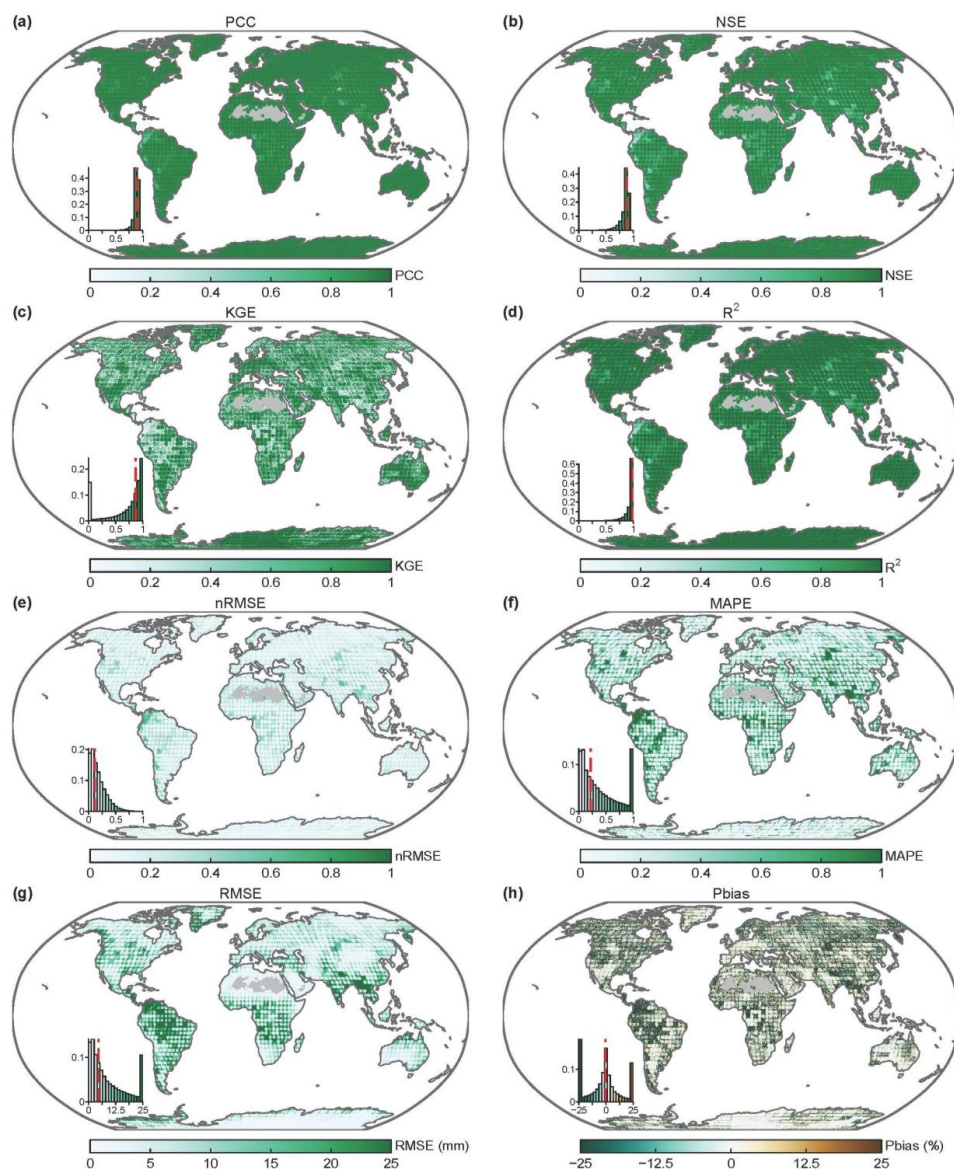
Figure 2. Performance of Random Forest in simulating JPL TWS anomalies under scheme 8 during the test period. Insets in each figure show the histogram of these metrics, with the vertical red line showing the median value. Data-sparse areas without reconstruction are marked in grey.

4 Performance evaluation of the GTWS-MLrec datasets

4.1 Comparison with GRACE/GRACE-FO observations

The final reconstructed long-term GTWS-MLrec datasets under the best-performing model are compared against GRACE/GRACE-FO observations. Unlike the comparison in Section 3.3 focusing on the test period, this section evaluates

the performance of GTWS-MLrec datasets in simulating the observations. The PCC, NSE, KGE and R^2 are larger than 0.9 in most land areas of the globe, and the nRMSE, MAPE, RMSE and Pbias also suggest low biases in most land areas (Figure 3; Figures S2-S3). We compare the values of these evaluation metrics with the scores achieved from previous reconstruction datasets (i.e., GRACE-REC and GRL-REC), which are also evaluated against the corresponding training GRACE/GRACE-FO solutions (Figures S4-S5). Across all the metrics, GTWS-MLrec typically achieves better scores than the previous two datasets. GTWS-MLrec, in particular, shows substantially lower values of nRMSE, MAPE, RMSE and Pbias than GRACE-REC and GRL-REC. It is informative to emphasize that unlike the GRACE-REC and GRL-REC datasets, which only reconstruct detrended and/or de-seasonalized TWS components, the three products of GTWS-MLrec provide the total components of TWS like GRACE/GRACE-FO observations. Therefore, it is not surprising that this reconstruction dataset shows such better performance when comparing with the observations.

The global yearly maps of reconstructed TWS fields are shown during the most recent strong El Niño years, i.e., 2015-2016. We find that GTWS-MLrec is able to capture the water storage anomalies due to the strong El Niño events. For example, eastern South America witnessed strong water deficits while most regions of South America suffered from severe flooding events during the onset period of El Niño in 2015 (Figure S6). By the wakening period of El Niño events in 2016, western and northern South America suddenly switched from pluvial floods to water deficit conditions (Figure S7). El Nino-Southern Oscillation (ENSO) is one of the leading climate oscillations, and often leads to extreme hydrological hazards over tropical regions (Juan et al., 2016); therefore, the anomalies in TWS over South America could be related to the strong El Niño events from the hydrological perspective (Emerton et al., 2017; Li et al., 2021). From the GRACE/GRACE-FO observations and our reconstructions, we also find that the 2015/2016 El Niño brought strong pluvial floods in southern China and most areas of Australia, and brought droughts to India, Middle East, central Europe and western America. The yearly TWSA map of the previous two reconstruction datasets (i.e., GRACE-REC and GRL-REC) are also depicted. We find that these two previous datasets also captured the El Niño event, but the GRACE-REC dataset failed to capture the fluvial signal in some regions such as southern Australia. In addition, GTWS-MLrec can also captures water deficit or wetness conditions in other strong El Niño events such as 1983 and 1998 (Figure S8). Overall, all the above results suggest that our reconstructions reliably characterize the anomalies induced by strong El Niño events, which is in line with the GRACE/GRACE-FO measurements.
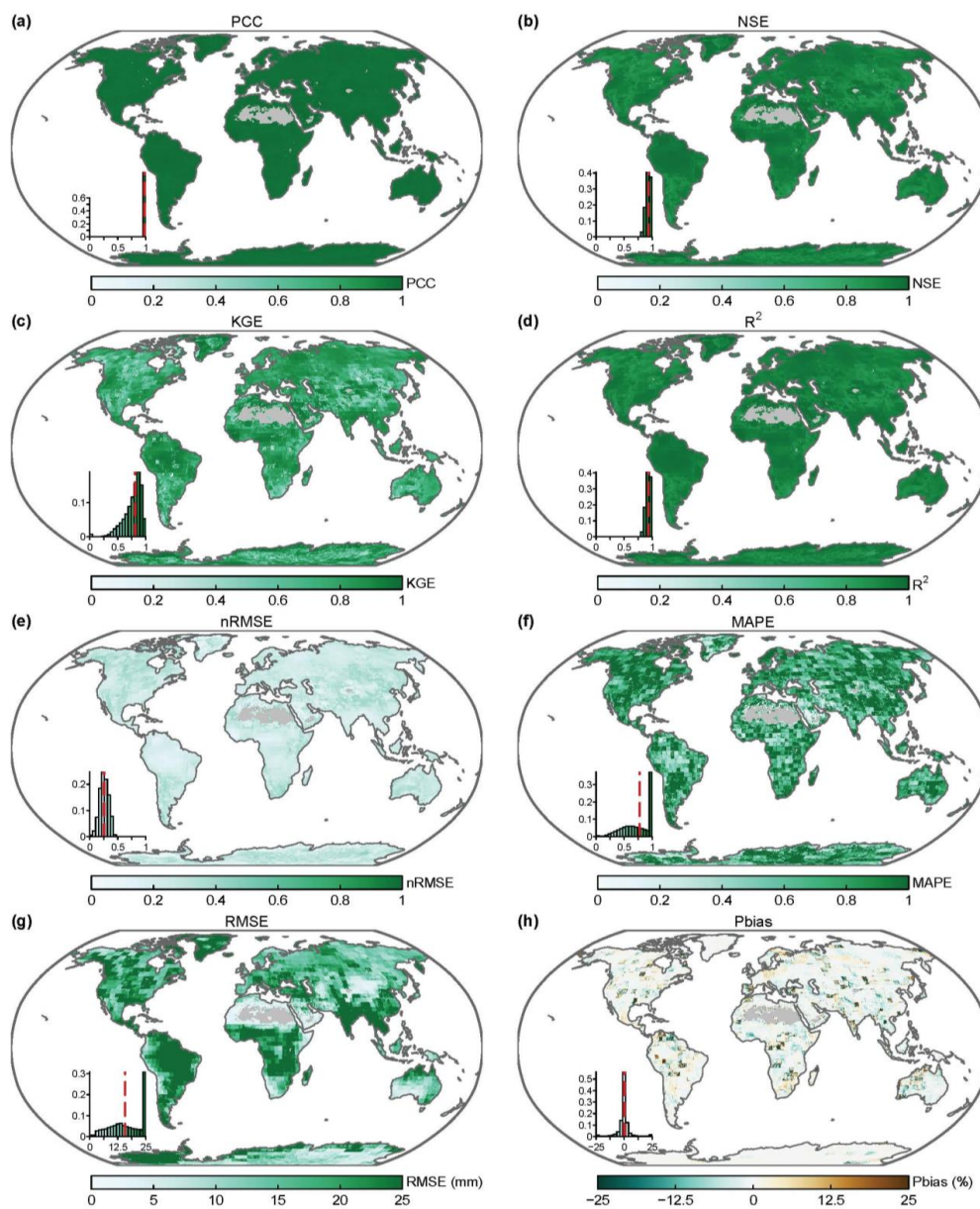
Figure 3 Comparison of our JPL-based GTWS-MLrec reconstructions against the GRACE/GRACE-FO observations. Insets
in each figure show the histograms of that evaluation metric, with the vertical red line showing the median value. Data-
sparse areas without reconstruction are marked in grey.

335

To evaluate the performance of GTWS-MLrec in capturing the global average TWS anomalies, we also present the long-
term monthly TWS anomalies over global land areas with and without Greenland and Antarctica (Figure 4). During the

GRACE era (2002-2022), the reconstructed TWS products do capture the inter-annual and seasonal cycle of GRACE/GRACE-FO observations well. Our three GTWS-MLrec products show similar long-term trends, which all indicate

340   a decreasing change in global average TWS anomalies. By comparing GTWS-MLrec with the GRACE-REC and GRL-REC datasets, we find that the reconstruction captures the inter-annual pattern of GRACE-REC well, while the GRL-REC dataset shows a slight overestimation phenomenon relative to our reconstruction and the GRACE-REC. Overall, the evaluations between GTWS-MLrec and GRACE/GRACE-FO observations indicates that our products achieve a good simulation performance at both grid and global scale.
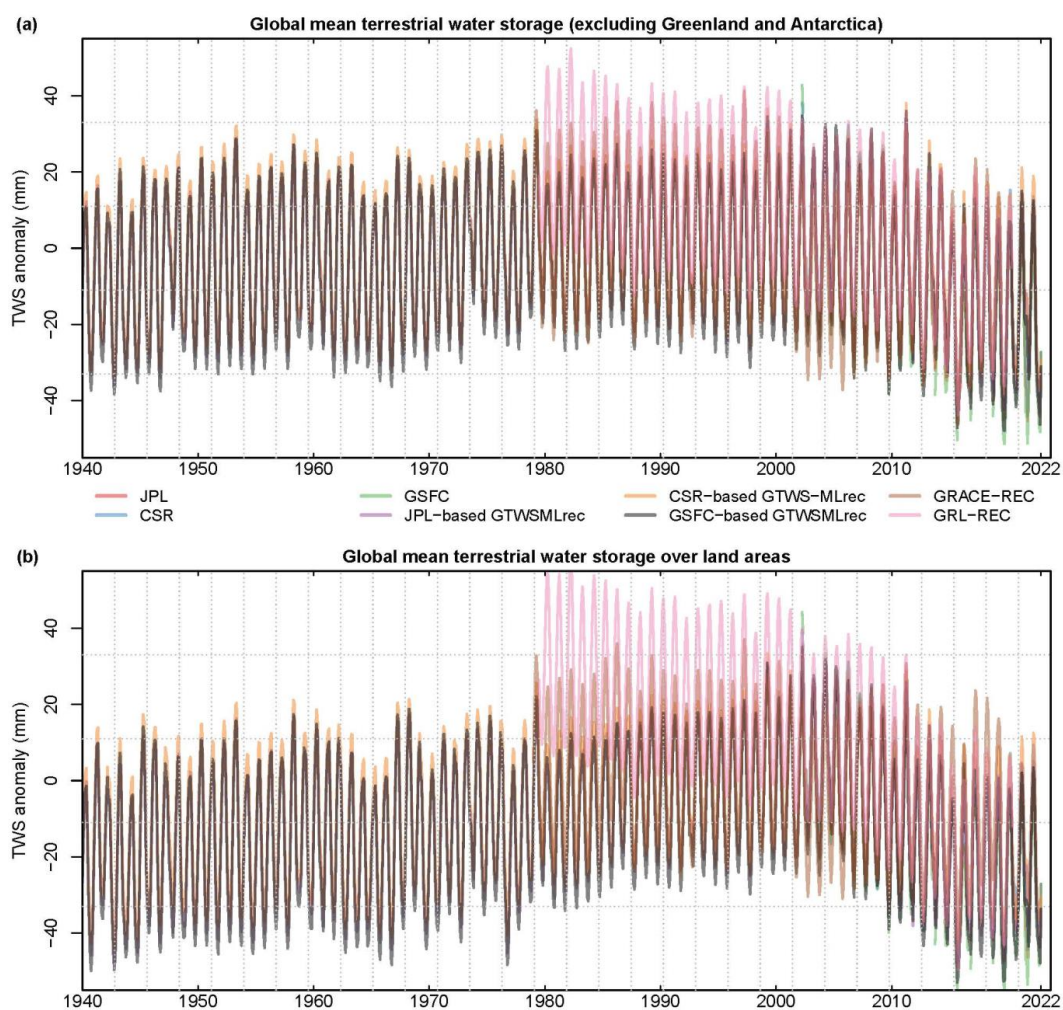
Figure 4. Global mean monthly terrestrial water storage anomaly derived by eight different datasets (including GRACE/GRACE-FO observations). (a) Global average TWS anomaly weighted by land area excluding Greenland and Antarctica; (b) Global average TWS anomaly over land areas.

4.2 Performance evaluation based on ocean-land water budget

Although the atmosphere holds moisture and thus can contribute to ocean mass changes by precipitation/evapotranspiration alternations, variations in ocean mass usually coincide with a comparable and opposite change of land water storage at a relatively long temporal scale (Chambers et al., 2016; Seo et al., 2023). Therefore, the sea level-based land-ocean mass budget provides an independent way to evaluate estimates of global average TWS variability. Here we assess the capacity of observed and reconstructed TWS products (GRACE/GRACE-FO, GTWS-MLrec, GRACE-REC, and GRL-REC) to constrain the sea level budget. The GMSL is employed to present ocean mass changes after subtracting the steric thermosteric and halosteric sea level. From the sea level budget-based analysis, we derive the time series of de-seasonalized and de-trended changes in ocean mass, which we compare against converted global average TWS estimates with consideration of land/ocean area (estimated after removing the trend and seasonal signal from the primary datasets). To ensure a consistent comparison among all candidate products, the Greenland and Antarctica are excluded when calculating the global average land TWS series. The comparison results show that, although all candidate datasets are well correlated with the ocean mass-based water budget, GTWS-MLrec and the GRL-REC product exhibit the strongest correlation with GMSL, with PCC values of >0.7 during 2006-2020 (Figure 5). Surprisingly, the reconstruction datasets also yield better results than the original GRACE/GRACE-FO datasets. This phenomenon has also been reported in previous studies (Humphrey and Gudmundsson, 2019), and probably because the global average GRACE/GRACE-FO TWS is more susceptible to non-compensating continental-scale errors (e.g., caused by biases from residual longitudinal stripes) compared to the data-driven reconstructions, which achieve smoother global average series. To evaluate our independent detrended and de-seasonalized GTWS-MLrec datasets, we also compare their converted TWS series with the GMSL. This converted dataset is different from directly excluding the trend and seasonal cycles from the full-component reconstruction datasets. The three detrended and depersonalized GTWS-MLrec datasets also show a good agreement with the global ocean mass changes, with a PCC ranging from 0.61 to 0.66 (Figure 6). All the results suggest that both of our reconstruction types are able to constrain the ocean-land water budget.
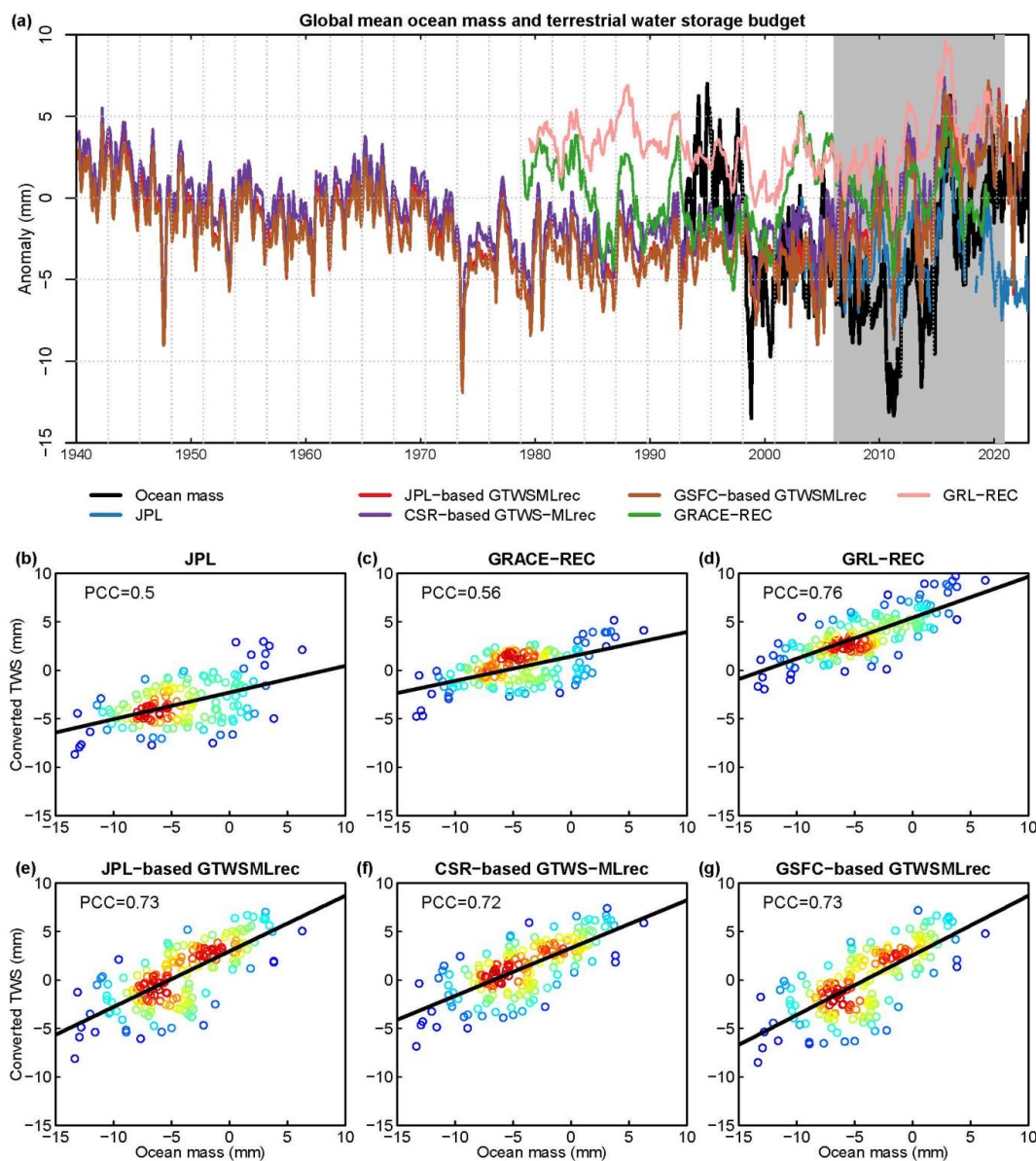
Figure 5. Comparison of the global average TWS variations (converted to equivalent sea level; unit: mm) against ocean mass changes derived from the sea level budget. **a**, the temporal dynamics of global average sea level and TWS. **b-g**, the regression plots of sea level and converted TWS; warmer (colder) colour indicates higher (lower) density of points; the black line denotes the linear regression. The converted TWS is estimated after extracting the trend and seasonal signals from the negative values of primary datasets by considering land-ocean area ratio.
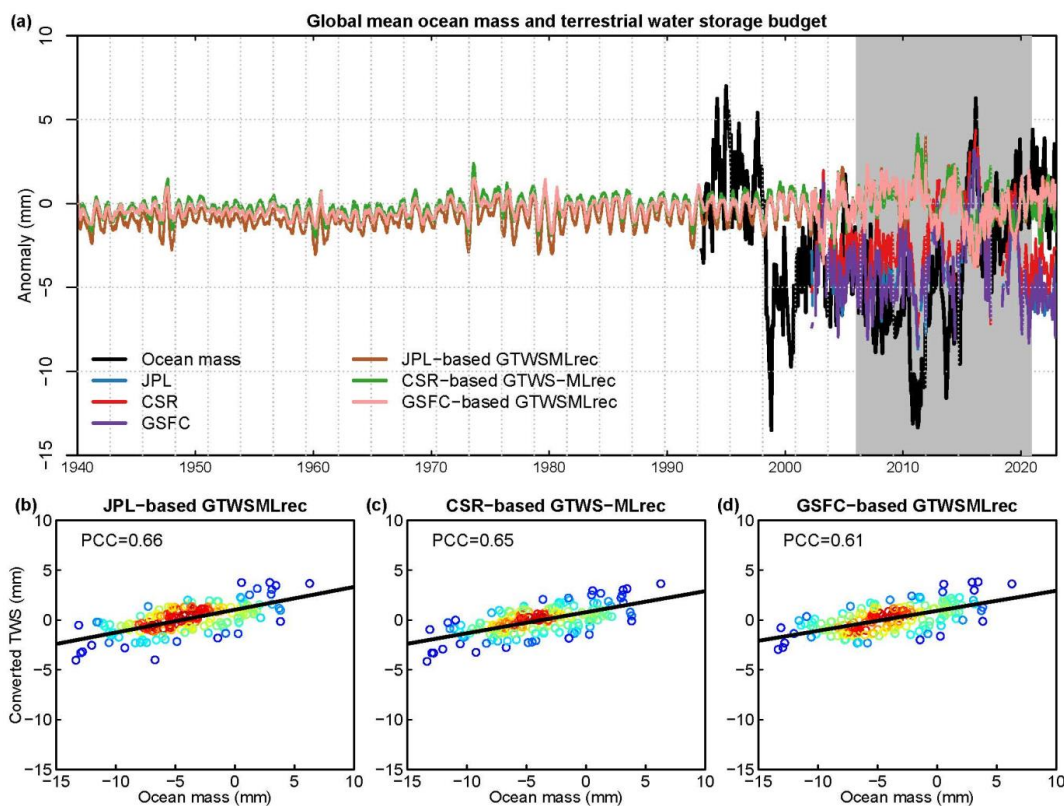
Figure 6. Global land-ocean mass budget based on our reconstructed detrended and de-seasonalized GTWS-MLrec datasets. **a**, temporal dynamics of global average sea level and TWS. **b-d**, regression plots of sea level and converted TWS; warmer
385 (colder) colour indicates higher (lower) density of points; the black line denotes the linear regression. The converted TWS is estimated as the negative values of primary detrended and de-seasonalized reconstruction datasets by considering land-ocean area ratio.

4.3 Performance evaluation based on water balance over large river basins

390 It is challenging to directly evaluate the performance of the TWS reconstructions prior to 2002 due to lack of GRACE/GRACE-FO observations. However, as TWS is a key element linking the atmospheric and terrestrial water balance budget, it is possible to independently evaluate the reconstructions by combining land and atmospheric moisture fluxes (Oki et al., 1995; Yin et al., 2022a). Hirschi and Seneviratne (2017) presented the BSWB dataset, which assimilated *in situ* streamflow observations and reanalysis-based moisture convergence in the atmosphere, and provided monthly variations in

395 TWS during 1979-2015 in 341 moderately large river basins (>100,000 km²) covering a variety of climate conditions. To provide a more comprehensive evaluation of our reconstructions, we extracted the variations in TWS from our reconstructions over the same basins as the BSWB dataset. As a caveat, we note that the BSWB is derived from the

predecessor of the ERA5 dataset (i.e., ERA-Interim), which is not entirely consistent with the main drivers of GTWS-MLrec products. Overall, both the GRACE/GRACE-FO and the reconstructed datasets do agree with the BSWB-based TWS

400   variations well (Figure 7a-h). For the JPL-, CSR- and GSFC- based mascons, these products show a PCC with BSWB ranging from 0.4 to 0.7 over most basins, and the global average PCC is about 0.58. For GTWS-MLrec, the three reconstructed datasets show a relatively similar pattern of PCC over the moderately large global river basins, further verifying the robust performance of our reconstructions.

405   We also compare the BSWB-based TWS variations with the GRACE-REC and GRL-REC products, and find a slightly higher PCC of GRACE-REC and GRL-REC than GTWS-MLrec reconstructions. This phenomenon might be due to two reasons. First, our GTWS-MLrec is reconstructed by assimilating the latest ERA5 dataset, which has substantial updates relative to ERA-Interim (i.e., inputs of BSWB). Second, the BSWB and two previous reconstruction datasets both mainly focused on climate-driven changes in TWS, but neglected underlying surface condition changes such as vegetation greening

410   due to increases in LAI and human-made infrastructures such as reservoirs. GTWS-MLrec not only assimilates climate variables but also captures changes in underlying surface conditions; therefore, these results may cause biases when comparing with the BSWB dataset. For example, the basins in China's Yangtze River usually show a higher PCC of previous reconstruction datasets than GTWS-MLrec. The Yangtze River has experienced rapid urbanization and reservoir constructions in recent decades (Gu et al., 2019), where the BSWB and previous reconstruction datasets have a similar TWS

415   pattern. We also compare the TWS variations of uncorrected and drift-corrected series from the BSWB with GTWS-MLrec, and find that the drift-corrected version of BSWB shows a slightly higher PCC with GTWS-MLrec (Figure 7i-j), suggesting the temporal smoothing improves TWS estimates by constraining the basin-scale water budget. Overall, our GTWS-MLrec achieves a relatively good agreement with the BSWB-based TWS variations in most large basins over most large basins of the globe.

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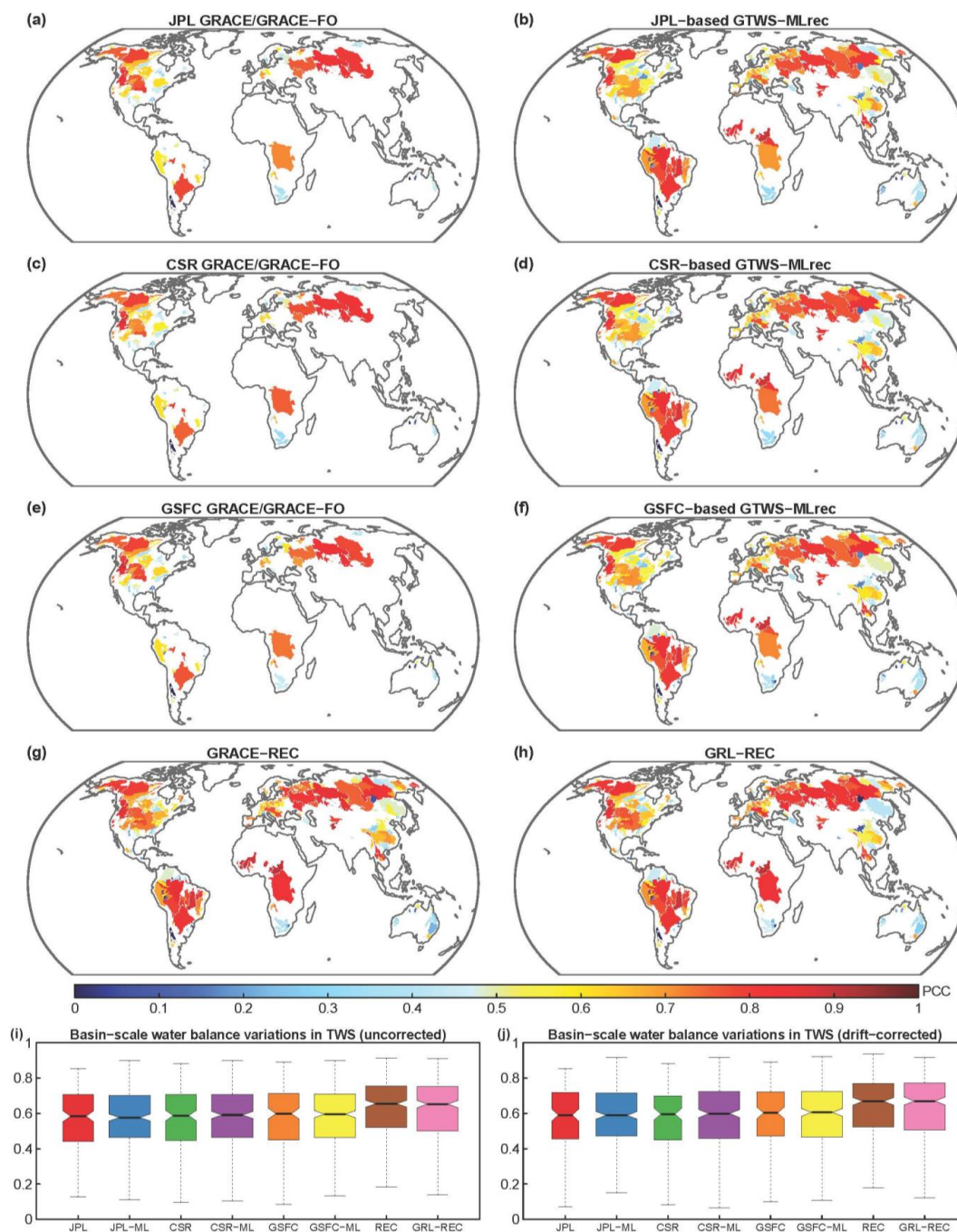
Figure 7. Comparison between variations in TWS derived from atmospheric basin-scale water balance (BSWB) dataset, GRACE/GRACE-FO measurements and reconstructions. **a-h**, PCC of drift-corrected BSWB-based TWS and observation/reconstruction datasets. **i-j**, boxplots of global PCC between BSCB and different TWS datasets for uncorrected

(**i**) and drift-corrected (**j**) BSWB datasets. In **i-j**, the JPL (or CSR, GSFC) denotes GRACE/GRACE-FO observations, and

425   the JPL-ML (or CSR-ML, GSFC-ML) denotes GTWS-MLrec reconstructions; the REC denotes GRACE-REC.

4.4 Comparison with annual streamflow measurements

As the BSWB dataset only provides TWS variations over large basins after 1979, we obtained streamflow data extending to

1940 from multiple sources, including the GRDC, USGS and the MWRC. After strict data quality control and screening of

430   the data records from 22538 stations, we retained 10168 hydrological stations in basins smaller than 100, 000 km^2 with at

least 10-year complete monthly streamflow series (see Section 2.3). River streamflow and TWS of course represent different

hydrological elements with divergent units; however, their temporal dynamics at a longer time scales (i.e., yearly) might be

correlated because runoff is an important water flux component in the total water storage (Rodell and Li, 2023; Yin et al.,

2023b). In addition, drier (or wetter) river streamflow state at an annual scale is usually related with anomalies in large-scale

435   atmospheric circulation, which is also represented by a similar signal indicated by annual TWS variations (Kang et al., 2023;

Yin et al., 2023a). To provide a better comparison with the basin-scale streamflow, the TWS from different reconstruction

datasets were aggregated at the watershed scale to obtain the catchment mean annual series in each basin using the Thiessen

polygon method. The mismatch in resolution between large-scale mass changes and local basin streamflow dynamics can

be partly alleviated by the spatial coherence of annual anomalies in weather/climate patterns. First, we compare the basin-

440   scale TWS against streamflow during 1979-2022, and find that GTWS-MLrec agrees well with the observational streamflow

dynamics at a yearly scale (Figure 8). The JPL- and GSFC- based GTWS-MLrec datasets both show a global average PCC

of 0.58, while the CSR-based reconstruction achieves a higher global average PCC of 0.60. The slightly better performance

of the CSR-based GTWS-MLrec product may be due to the fact that the primary CSR mascon solution has a finer spatial

resolution than the other two products. The previous GRACE-REC and GRL-REC reconstruction datasets show poorer

445   performance than GTWS-MLrec products in terms of the PCC. The GRACE-REC and GRL-REC datasets show a global

average PCC of 0.56 and 0.47, respectively. These results suggest that GTWS-MLrec achieves the best performance relative

to streamflow, in terms of reproducing past water cycle variability. In addition, the streamflow and TWS from our three

reconstructions are compared during 1940-1980 (Figure 9). Our GTWS-MLrec products show a global average PCC of 0.55-

0.57, and the CSR-based dataset still shows the best performance, suggesting a good agreement with the temporal streamflow

450   dynamics.
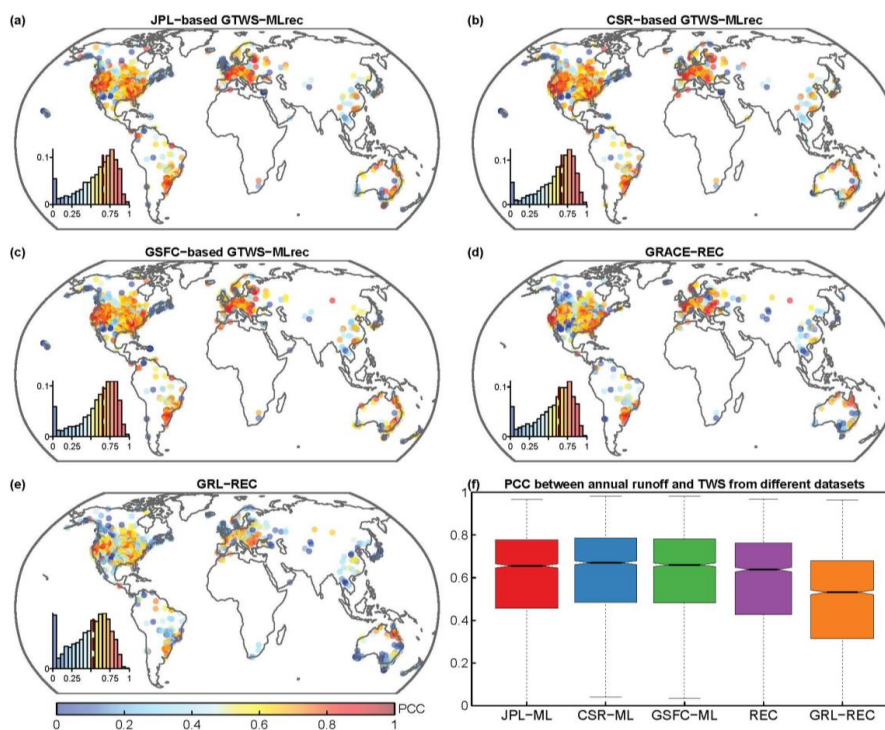
Figure 8. Correlation of annual streamflow and aggregated basin-scale TWS from the different reconstruction datasets during 1979-2022. **a-e**, Global distribution of PCC for the different datasets; **f**, Boxplot of the PCC for all stations globally; the REC denotes GRACE-REC. Insets in each figure show the histogram of these metrics, and the vertical red line shows the median value.
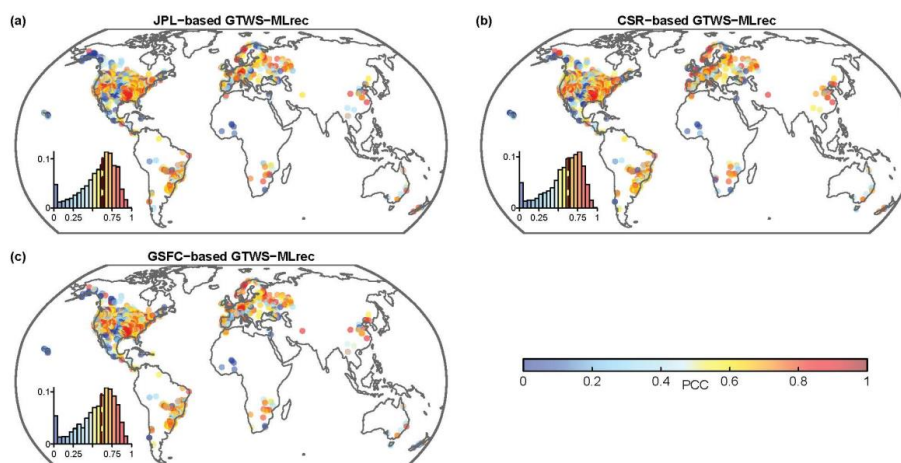
Figure 9. Correlation of annual streamflow and aggregated basin-scale TWS from different reconstruction datasets during 1940-1980. Insets in each figure show the histogram of these metrics, and the vertical red line shows the median value.

5 Data availability

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The GTWS-MLrec dataset is archived on Zenedo at the link: https://zenodo.org/record/8187432 (Yin et al., 2023c). It is distributed with a CC-BY license. The uploaded data provided are (a) JPL-based GTWS-MLrec TWS series; (b) CSR-based GTWS-MLrec TWS series; (c) GSFC-based GTWS-MLrec TWS series; (d) JPL-based GTWS-MLrec TWS constructions after removing trend and seasonal signals; (e) CSR-based GTWS-MLrec TWS series after removing trend and seasonal

465 signals; (f) GSFC-based GTWS-MLrec TWS series after removing trend and seasonal signals; (g) Global average TWS series over land areas; (h) Global average TWS series over land areas after excluding Greenland and Antarctica.

6 Summary, applications, and outlook

TWS—which encompasses all water storage or fluxes on the land, such as groundwater, soil moisture, snow/ice cover, and

470 surface water—is a key element affecting the hydrological cycle at both global and regional scales. TWS also plays an essential role in the global water and energy budget balance, which is highly correlated with underlying surface conditions (e.g., vegetation) and climate fluctuations (e.g., El Niño events). Thus, to improve our understanding of changes in the hydrological cycle under climate change, a long-term TWS dataset is urgently needed. However, direct remote sensing-based observations of TWS from the GRACE/GRACE-FO only go back to 2002. Therefore, recent works have

475 retrospectively reconstructed TWS at global or regional scales (e.g., Humphrey and Gudmundsson 2019; Sun et al., 2020; Li et al., 2021, 2022). Despite being useful for various purposes, these studies have been usually constrained by the short time period (e.g., starting from 1979) or coarse spatial resolution ($\geq$ 0.5°). Furthermore, previous studies have often focused primarily on climate-driven changes in TWS and did not fully assimilate information about vegetation conditions such as LAI. GTWS-MLrec, the new TWS product presented here, provides a monthly TWS dataset from 1940 to the present at

480 0.25°, and is reconstructed by assimilating a large number of meteorological, hydrological, human-relevant and vegetation variables. By comparing GTWS-MLrec against GRACE/GRACE-FO observations, we find that GTWS-MLrec usually achieves a higher correlation coefficient and lower biases than previous reconstruction datasets. We also evaluate the performance of different reconstruction datasets by using the land-ocean mass budget, basin-scale water balance and temporal streamflow dynamics as well as their ability to capture El Niño events. These independent evaluations all suggest

485 that our GTWS-MLrec achieves better performance than previous datasets under most conditions.

We envision that GTWS-MLrec, with its comprehensive and extensive attributes, could provide rich information as references across a broad range of geoscience-relevant applications. First, the long-term TWS series can provide rich information for understanding changes in the global and regional hydrological cycle under climate change. For example, the

490   TWS-based drought index has shown great potential for monitoring and assessing large-scale droughts in numerous regions of the globe (Long et al., 2014; Yin et al., 2022a). Second, the TWS anomalies of GTWS-MLrec could be used as an independent benchmark to evaluate the performance of a set of hydrological/climate models. Numerous studies have compared the TWS produced by GRACE/GRACE-FO and hydrological models, and have found that many physical models tend to simulate an earlier peak of TWS in seasonal cycle, which may be related to an underestimation of the overall water

495   storage capacity (Schellekens et al. 2017; Long et al., 2017). However, these comparisons were conducted over a short time period, while GTWS-MLrec will be helpful to achieve a more robust long-term assessment. Third, GTWS-MLrec will provide more useful information for understanding terrestrial carbon-climate feedbacks. As the TWS reflects the water condition needed for vegetation growth, TWS variability is tightly coupled to the atmospheric CO_2 growth rate (Liu et al., 2023). Therefore, our GTWS-MLrec could contribute to understanding the possible shifts in terrestrial climate-carbon

500   coupling under climate change. Finally, GTWS-MLrec can also improve the understanding of the dynamic relationship between climate change and human activities, such as groundwater depletion (Long et al., 2020; Seo et al., 2023), irrigation (Lv et al., 2019), urbanization (Huang et al., 2023), compound hazards (Yin et al., 2022b), and reservoir operation (Shah et al., 2019).

505   Although GTWS-MLrec presents substantial improvements over previous TWS reconstructions and holds potential for a broad range of geoscience applications, a few caveats should be acknowledged. The main inputs for feeding the machine learning model are sourced from the ERA5 dataset, which provides hourly and monthly climatic variables in 0.25°. Many basin-scale hydrological studies may need a TWS dataset at a finer spatial resolution, where GTWS-MLrec may be too coarse to constrain the water budget balance. The ERA5-Land dataset is available 0.1° resolution, but is not considered in

510   this study because it lacks some key variables such as cloud cover and moisture divergence. The estimation accuracy of the TWS reconstruction relies on the quality of the ERA5 datasets, and next generations/versions of the TWS data could potentially be improved as ERA5 itself improves over time. Although we have tried to assimilate information of underlying surface condition and vegetation states, the series of SIF and ISA cannot be extended to 1940 because they are constrained by the availability of historical satellite records. To address this issue, we have designed eight different data schemes in the

515   machine learning models, and the long-term LAI series can also provide information about vegetation. The human-made hydraulic infrastructures such as reservoir may play an important role in regulating TWS, but these effects have not been fully considered in the GTWS-MLrec. Future works may seek to incorporate more predictors in the long-term TWS reconstruction and to provide more systematic independent evaluations.

520 **Author contributions**. J.Y. conceived the study, performed the analyses, and wrote the paper. L.S., P.L., F.L., Y.P., and P.G. provided critical input and assisted in interpretation of the results. A.K. and L.Y. contributed to data analysis by Google Earth Engine. All the co-authors reviewed and edited the paper.

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685