



1	A coarse pixel scale ground "truth" dataset based on the
2	global in situ site measurements to support validation and
3	bias correction of satellite surface albedo products
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12	Abstract- In situ measurements from sparsely distributed networks worldwide are
13	a valuable source of the reference data for validating or correcting the bias of satellite
14	products. However, the significant differences in spatial scale between in situ and
15	satellite measurements make them incomparable except for that the underlying surface
16	of in situ sites is absolutely homogeneous. Instead, in site measurements need to be
17	upscaled to be matched with satellite pixel. Based on the upscaling model we proposed
18	as well as the consideration that in-situ observation generally lacks spatial
19	representativeness due to the widely distributed spatial heterogeneity, we have
20	developed a coarse pixel-scale ground "truth" dataset based on ground measurements
21	of 368 in situ sites from the sparsely distributed observation networks. Furthermore, the
22	effectiveness of the dataset was carefully evaluated over the sites with different degrees
23	of spatial representativeness. The results demonstrate that using this dataset in
24	validation outperforms the direct comparison between satellite and in situ site
25	measurements. The accuracy of the reference data employed for validation or bias
26	correction can be enhanced by 3.5% overall with this dataset. But the performance of
27	the dataset show dependence on the degree of spatial heterogeneity. Specifically, the
28	improvement of accuracy was the most significant over the regions with strong spatial
29	heterogeneity, with the accuracy of reference data enhanced by 7.3% . To the best of our
30	knowledge, this dataset is unique in providing coarse pixel scale ground truth with the





31 widest spatial distribution and longest time series. Its ability to capture both spatial and

32 temporal variations of surface albedo at coarse spatial scales makes it an invaluable

33 resource for validating and correcting global surface albedo products.

1. Introduction

Surface albedo is an important variable in climate and biogeochemical models 35 because it determines the amount of energy absorbed by the earth surface. Coarse-pixel 36 (i.e., with a km pixel scale) satellite albedo products such as MODIS and NPP VIIRS 37 have been widely used to tackle with global challenges and support a range of initiatives 38 (e.g., The Paris Agreement and Sustainable Development Goals). However, satellite 39 albedo products generally suffer from different degrees of errors due to the error of 40 41 satellite observation data and the limitation of inversion algorithm, and the error of remote sensing product brings great uncertainty to the next application of the product. 42 Taking albedo as an example, the change of albedo of 11% will cause a fluctuation of 43 surface net radiation of 3.5 Wm-2 on global and annual averages (GCOS, 2011), which 44 in turn will cause the change in global temperature of 0.1K. An increase of $0.00106\pm$ 45 46 0.00008 (mean \pm standard deviation) of albedo will cause the radiation at the top of the atmosphere to cool by -0.15 ± 0.1 Wm-2 (Ghimire et al., 2014). Therefore, it is very 47 important to evaluate the uncertainty of remote sensing albedo products. In particular, 48 when the error is relatively large, it is urgent to correct the error of remote sensing 49 albedo products to improve the application accuracy of remote sensing products. 50

Both the validation and correction of remote sensing products depend on the 51 reference data which can represent the ground truth on the coarse pixel scale. Ground 52 measurement is usually regarded as the truth of the target variable on the observation 53 54 scale. However, due to the scale mismatch between ground measurement and satellite pixel as well as the widely distributed spatial heterogeneity, ground measurement can 55 56 not be directly used as the absolute truth on the pixel scale (Wu et al., 2019). Limited by the means and methods of ground measurement, the absolute truth on the coarse 57 pixel scale can not be obtained. Instead, its best approximation value can be obtained. 58 What we can do is to get the pixel scale relative truth close to the absolute truth as 59 accurately as possible (Wu et al., 2019). For conciseness, the pixel scale relative truth 60 is described as pixel scale truth in this paper. 61

The sparsely distributed in situ sites (i.e., at most one site within a specific product grid cell) from the networks such as FLUXNET, BSRN, and SURFRAD provide an important data source for the validation of remote sensing albedo products.(Chu et al.,





2021; Augustine et al., 2000; Driemel et al., 2018). However, in situ measurements 65 cannot be directly used as the coarse pixel scale truth given that the footprint of in situ 66 sites is far less than the scale of a coarse pixel. A practical method of using in situ site 67 measurements as the coarse pixel scale truth is to conduct the spatial representativeness 68 assessment of in situ sites (Román et al., 2009; Wang et al., 2014; Moustafa et al., 2017). 69 70 However, since these in situ sites were not originally established for the validation or 71 bias correction of satellite products, only a small part of them was proved to be spatially representative, and most of them were rejected. Even for the representative site, the 72 representativeness errors of in situ measurements are still inevitable, because land 73 surface are not absolutely homogeneous throughout the year (Colliander et al., 2017; 74 Xu et al., 2018; Lei et al., 2018; Williamson et al., 2018). Consequently, the 75 representative in situ measurements are only limited to a few locations on the globe and 76 cover discrete time periods, which cannot support a comprehensive validation and bias 77 correction over a wide range of conditions (Loew et al., 2016). 78

To overcome the representative errors of in situ measurements and promote 79 80 utilization ratio of in situ sites from these sparse networks in validation, Wu et al. (2020) have proposed an upscaling method specified for the single site in situ measurements. 81 However, the effectiveness of this method has not been comprehensively assessed and 82 its transferability to in situ sites all over the world is still unknown. In this study, we 83 first evaluated the performance of the upscaling model over the 368 in situ sites 84 85 throughout the world. Then the pixel scale ground truth data with a spatial resolution of 0.5 km were produced based on these in situ measurements. Finally, the potential 86 usage of this dataset was discussed. As far as we know, this dataset is unique in 87 providing coarse pixel scale ground truth with the widest spatial distribution and longest 88 time series. 89

90 2. The experimental data

91 2.1 In situ site observation

In this study, in situ sites from Surface Radiation (SURFRAD), Baseline Surface Radiation Network (BSRN), AmeriFlux, EuropFlux, HiwaterWSN (Che et al., 2019), and Huailai station (Li et al., 2016; Ma et al., 2013), were utilized to generate the coarse pixel scale "truth" dataset. The spatial distribution of these in situ sites is displayed in Fig. 1. It can be seen that these sites are unevenly distributed, with most of them concentrated in the Northern Hemisphere. With consideration of the completeness and





- 98 the quality of in situ observations, a total of 368 sites were selected in this study, which
- 99 covers the main vegetation functional types and climatic regions (Fig. 2). The wide
- 100





Figure 1. Global distribution of the 368 sites over different land cover types.



Figure 2. The distribution of stations across fifteen Land cover types indicated by International Geosphere-Biosphere Programme (IGBP). WAT, BSV, SNO, CVM, URB, CRO, WET, GRA, SAV, WSA, OSH, CSH, MF, DBF, EBF, and ENF are the abbreviation of water bodies, Barren, snow and ice, cropland/natural vegetation mosaics, urban and built-up lands, cropland permanent wetlands, grassland, savanna, woody savanna, open shrubland, closed shrublands, mixed forest, deciduous broadleaf forest, evergreen broadleaf forest, and evergreen needleleaf forests. The numeric value displayed above each bar in the chart indicates the total number of stations associated with the corresponding land cover type.





These in situ sites were equipped with two pyranometers mounted back-to-back, 111 one pointed downward and the other upward, measuring the downward radiation and 112 upward radiation (Song et al., 2019). It is important to note that the footprint of in situ 113 sites is not fixed, given that the measurement height of the albedometers from the 114 underlying surface varies from site to site (Marion, 2021; Song et al., 2019). These 115 116 radiometers have been rigorously calibrated and continuously supervised to reduce systematic measurement errors(Jia et al., 2013; Wang et al., 2009; Zhou et al., 2016). 117 In situ site measured albedo of the local solar noon was calculated using the ratio of the 118 mean upward radiation to the mean downward radiation between 11:00 a.m. and 1:00 119 p.m. local time as suggested by Lin et al. (2022). 120

121 2.2 The high-resolution albedo

The high-resolution albedo data were used in this study for two purposes. One is 122 123 to calculate the upscaling coefficients for in situ measurements since they can capture the spatial variation characteristics of surface albedo within the coarse pixel extent. The 124 other is to serve as the reference data because they can represent the coarse pixel albedo 125 after aggregation. Landsat7 ETM+ images were used to generate high-resolution albedo 126 due to their high spatial resolution of 30 m. This satellite has a revisit period of 16 days 127 128 and an inclination of 98.2°, making it possible to acquire data up to a latitude of 82° N 129 (Rösel et al., 2011).

130 The Landsat Level-2 reflectance data, which have been atmospherically and geometrically corrected (Teixeira et al., 2020), were used in this study. Furthermore, 131 additional processing such as cloud screening (with Fmask algorithm) and cloud 132 removal (with Cloud-removal-by-Synthesis (CRS) algorithm) was also carried out to 133 further enhance the quality of the surface reflectance data. The retrieval of high-134 resolution surface albedo was based on the algorithm proposed by Liang et al. (2001), 135 which provided narrowband-to-broadband conversion coefficients (Eq. (1)) for albedo 136 retrieval from Landsat7 ETM+ using spectral reflectance library and simulations under 137 138 different atmospheric and surface conditions under the Lambertian assumption. The anticipated accuracy of this algorithm approximates 0.02 (Liang, 2001). This equation 139 provides a valuable tool for the accurate estimation of surface albedo from satellite data 140 and has been widely used in numerous studies on land surface processes and global 141 climate change. 142

143 $\alpha_{short} = 0.356\alpha_1 + 0.130\alpha_3 + 0.373\alpha_4 + 0.085\alpha_5 + 0.072\alpha_7 - 0.0018$ (1) 144 where α_{short} denotes the shortwave surface albedo, and α_i denotes the spectral





albedo at the wavelength of *ith* satellite spectral band.

146 2.3 Coarse pixel-scale satellite albedo product

The MCD43A3 V061 product was used in this paper to serve as an example of coarse pixel satellite albedo products due to its wide acceptance. The shortwave (3000-5000 nm) albedo was extracted to match the spectral range of in situ-measured albedo. It provides black sky albedo (BSA) and white sky albedo (WSA) with a spatial resolution of 500 m and a temporal resolution of daily (Schaaf et al., 2002). The BSA and WSA can be linearly combined with the sky diffuse light ratio to derive the bluesky albedo in the actual environment (Pinty et al., 2005).

154 $\alpha = \alpha_{WSA} \times r + \alpha_{BSA} \times (1 - r)$ (2)

where α , α_{WSA} , and α_{BSA} represent the blue-sky albedo, WSA, and BSA of MCD43A3, respectively. The actual sky diffuse light ratio r can be calculated using equation (3) as suggested by Stokes and Schwartz (1994).

158 $r = 0.122 + 0.85 \times \exp(-4.8 \times \cos \theta)$ (3)

159 where θ denotes the solar zenith angle at local solar noon.

160 **2.4 The ancillary dataset**

The auxiliary data were used in this paper to explore the potential influence factors 161 162 on the accuracy of the pixel scale ground truth data. Several common surface parameters, including elevation, land cover types, and spatial heterogeneity, were 163 considered. DEM data were obtained from the Multi-Error-Removed Improved-Terrain 164 (MERIT) Digital Elevation Model (DEM) (Mcclean et al., 2020; Yamazaki et al., 2017), 165 with a high horizontal resolution of three arc-seconds (approximately 90 meters). The 166 MERIT DEM was produced using the advanced algorithms of HydroSHEDS 167 (Hydrological data and maps based on SHuttle Elevation Derivatives at multiple Scales) 168 to remove various types of errors in existing DEMs, such as voids, noise, and artifacts 169 (Uuemaa et al., 2020). The MCD12Q1 product provides global land cover type 170 171 information with a spatial resolution of 500 m at an annual time step from 2012 to 2018. It provides information on 17 different land cover types, making it a valuable data 172 source for many applications. It is important to point out that there is a slight difference 173 regarding the land cover information between the data of different years. Here, we have 174 used the land cover type with the highest degree of agreement between the data in 175 176 different years.





177 Spatial heterogeneity is a critical factor influencing the spatial scale match between in situ and satellite measurements because it reduces the spatial 178 representativeness of in situ measurements (Wu et al., 2022). It refers to the uneven 179 distribution of surface albedo within the coarse pixel. A pixel that exhibits spatial 180 heterogeneity denotes that the value of surface albedo at one location is different from 181 182 that of other locations. To quantify the spatial heterogeneity of surface albedo within a coarse pixel, the spatial variability (standard deviation, Std) of all subpixel albedos 183 within a coarse pixel was calculated (Colliander et al., 2017; Jin et al., 2003). Here, the 184 subpixel albedos denote the high-resolution pixel albedo within the coarse pixel. 185

$$Std = \sqrt{\frac{1}{L-1} \sum_{i=1}^{L} (Z_i - \overline{Z})^2}$$
(4)

187 where Z_i denotes the high-resolution albedo and \overline{Z} is the averaged albedo of all 188 high-resolution albedos within the extent of the coarse pixel. *L* refers to the number of 189 high-resolution albedo pixels with a coarse pixel.

190 **3. Methodology**

186

191 **3.1** The upscaling model specified for single in situ site measurements

192 In situ measurements taken at a single in situ site can provide accurate 193 measurements on the point scale and offer continuous temporal variation information for long time series. But they are insufficient to represent albedo at the coarse-pixel 194 195 scale due to the spatial heterogeneity within the coarse pixel. High-resolution albedo 196 maps can capture the spatial variation information within the coarse pixel. The basic idea of the upscaling model is to derive the upscaling coefficients based on high-197 198 resolution albedo maps and then applied these upscaling coefficients to long-term in 199 situ measurements (Wu et al., 2020). In this way, both the spatial variation information 200 and the temporal variation information of surface albedo can be captured through the combination of high-resolution albedo maps and long time series in situ measurements, 201 deriving the long time series pixel scale ground truth data. 202

Since the high-resolution albedo maps serve as an important linkage between in situ measurement scale and satellite coarse pixel scale, they should meet several acquirements. First, its spatial resolution should be equal or close to the footprint of in situ measurements, because the upscaling coefficients were determined by long time series high-resolution albedo maps and then were applied to in situ measurements. Second, the high-resolution albedo maps should cover at least one full cycle period,





(4)

which is typically one year, to account for the seasonal variations of surface heterogeneity caused by phenology, because the upscaling coefficients will be applied to long time series in situ measurements. In this study, the Landsat7 ETM+ albedo maps from 2012 to 2018 were processed to capture the spatiotemporal variation of surface albedo and used as prior knowledge for deriving the upscaling coefficients.

The upscaling coeffcients were calculated for each subpixel within the coarse pixel extent by establishing a regression relationship between one subpixel albedo time series and the subpixel albedo time series corresponding to the in situ site (Eq. (4)). To avoid the uncertainty caused by different data sources, both of them were simulated by Landsat7 albedo. Using the same data source can reduce the influence of errors of Landsat7 ETM+ albedo to a certain extent.

220
$$\theta_{ETM+}(x, y, d) = W(x, y)^{T} \theta_{ETM+in \ situ}(d)$$

221 and

222

230

$$\theta_{ETM+\ in\ situ}(d) = [1, \theta_{ETM+\ in\ situ}(d)]^T$$
(5)

where x and y correspond to the location of a single Landsat7 ETM+ pixel within a coarse pixel, while d denotes the date of the Landsat7 ETM+ albedo map. $\theta_{\text{ETM+}}$ denotes the individual Landsat7 ETM+ pixel within a coarse pixel. $\theta_{\text{ETM+},\text{in situ}}$ indicates the Landsat7 ETM+ pixel albedo time series corresponding to the in situ site. The vector W represents the upscaling coefficients to be derived.

To ensure a robust estimation, a cost function J is established by combining all the Landsat7 ETM+ albedo data throughout the whole time series (i.e., 2012-2018).

$$J = \min\{\sum_{d=1}^{L} [\theta_{ETM+}(x, y, d) - W(x, y)^{T} \theta_{ETM+_{in} situ'}(d)]^{2}\}$$
(6)

Using the ordinary least-squares (OLS) algorithm, the vector of coefficients W canbe obtained by minimizing the cost function.

233
$$W(x,y) = \left[\theta_{ETM+_in\ situ}^{T}\theta_{ETM+_in\ situ}\right]^{-1}\theta_{ETM+_in\ situ}^{T}\theta_{ETM+_in\ situ}\theta_{ETM+_in\ situ}$$
(7)

After the upscaling coefficients were acquired, they were applied to in situ site measurements to simulate the in situ reporting of surface albedo over each Landsat7 ETM+ pixel ($\theta_{in \ situ \ ETM+}$) within the coarse pixel:

237 $\theta_{in \ situ \ ETM+} \ (x, y, d) = W(x, y)^T \theta_{in \ situ}(d) \tag{8}$

Then the coarse pixel ground truth ($\theta_{in \ situ_ref}$) can be derived by aggregating all the $\theta_{in \ situ \ ETM+}$ within the coarse pixel using the point spread function (PSF) of the MODIS albedo characterized by Peng et al. (2015).





$$\theta_{in \ situ_ref}(d) = \frac{\int_{(x,y)\in D} f_{PSF}(x,y)\theta_{in \ situ \ ETM+}(x,y,d)}{\int_{(x,y)\in D} f_{PSF}(x,y)}$$
(9)

where *D* denotes the spatial extent of the coarse pixel, and f_{PSF} represents the PSF. As shown in Eq. (8-9), it can be seen that the upscaling coefficients are timeindependent and can be applied to in situ measurements throughout the entire time series under the condition that there are no sudden changes such as wildfires and deforestation on the land surfaces. It is noteworthy that a high-resolution albedo map is no longer a prerequisite for the practical upscaling process once the upscaling coefficients have been obtained.

249

3.2 The evaluation of upscaling models and pixel scale ground truth

251 **3.2.1** The evaluation of the upscaling model

As explained above, the key to the upscaling method is to obtain upscaling 252 coefficients based on 30-meter Landsat7 ETM+ albedo data from 2012 to 2018. Hence, 253 the accuracy of the upscaling model is highly dependent on the performance of the 254 upscaling coefficients. Inspired by the evaluation method proposed by Wu et al. (2016), 255 the accuracy of upscaling coefficients were assessed using the Landsat7 ETM+ albedo 256 data from 2019 to 2021. The upscaling results with the upscaling coefficients can be 257 determined with the Eqs. (10-11). The reference value on the coarse pixel scale is the 258 aggregated Landsat7 ETM+ albedo at the 500 m resolution as recommended by Wu et 259 al. (2016) (Eq. (12)). 260

$$\theta_{RETM+}(x, y, d) = W(x, y)^T \theta_{ETM+}(d)$$
(10)

261 262

263
$$\theta_{upscaling}(d) = \frac{\int_{(x,y)\in D} f_{PSF}(x,y)\theta_{RETM+}(x,y,d)}{\int_{(x,y)\in D} f_{PSF}(x,y)}$$
(11)

264
$$\theta_{reference}(d) = \frac{\int_{(x,y)\in D} f_{PSF}(x,y)\theta_{ETM+}(x,y,d)}{\int_{(x,y)\in D} f_{PSF}(x,y)}$$
(12)

where θ_{RETM+} is the reporting of surface albedo over each Landsat7 ETM+ pixel based on the Landsat7 ETM+ pixel albedo corresponding to the in situ sites and the upscaling





coefficients. $\theta_{upscaling}$ denote the upscaling results and $\theta_{reference}$ represent the reference value. 267 The similarity and consistency between the $\theta_{upscaling}$ and $\theta_{reference}$ were evaluated 268 by three metrics: bias, coefficient of determination (R²), and root-mean-square error 269 (RMSE). 270

271
$$RMSE = \sqrt{\sum_{d=1}^{L} (\theta_{upscaling}(d) - \theta_{reference}(d))^2 / L}$$
(10)

272
$$Bias = \sum_{d=1}^{L} (\theta_{upscaling}(d) - \theta_{reference}(d))/L$$
(11)

273
$$\mathbf{R} = \frac{\sum_{d=1}^{L} (\theta_{upscaling}(d) - \overline{\theta}_{upscaling})(\theta_{reference}(d) - \overline{\theta}_{reference})}{[\sum_{d=1}^{L} (\theta_{upscaling}(d) - \overline{\theta}_{upscaling})^2 \sum_{d=1}^{L} (\theta_{reference}(d) - \overline{\theta}_{reference})^2]^{1/2}}$$
(12)

3.2.2 Assessment of pixel scale ground truth 274

The way of evaluating the pixel scale ground truth is similar to that of the upscaling 275 model. Namely, the accuracy of $\theta_{in situ ref}$ was assessed through the comparison with 276 $heta_{\textit{reference}}$. Furthermore, the performance of the pixel scale ground truth is also compared 277 to that of single in situ site measurements. To eliminate the influence of the magnitude 278 of surface albedo on accuracy indicators, the relative root-mean-square error (RRMSE) 279 was used here, which is defined as the ratio of RMSE to the mean value of surface 280 281 albedo on the coarse pixel scale. 282

$$RRMSE = \frac{RMSE}{\theta_{in \ sutu.ref}} \times 100\%$$
(13)

where $\overline{\theta_{in \ situ_{ref}}}$ represents the mean value of coarse pixel-scale albedo truth, and L 283 denotes the length of the temporal sequence of data. 284

4. Results and Discussion 285

286 4.1 The performance of the upscaling model

The performance of the upscaling coeffcieints has been comprehensively 287 evaluated over the 368 in situ sites. The wide spread of in situ sites across different 288 elevations, different land cover types, and different degrees of spatial heterogeneity can 289 ensure the objectivity of the evaluation results. To show the agreement between 290 291 $\theta_{upscaling}$ and $\theta_{reference}$ more intuitively, we present the scatterplots between them in Fig. 2. As shown in Fig. 3, the scatterplots between $\theta_{usscaling}$ and $\theta_{reference}$ are generally 292 distributed around the 1:1 line, with R² close to 0.9. The upscaling coefficients show no 293





systematic error, indicated by the biases close to 0. However, the performance of the upscaling models is site dependent. For instance, the accuracy of the upscaling results over the US-Ha2 site obviously outperforms that of the IT-Tor site (Fig. 2(c vs. f)). To fully understand the performance of upscaling coefficients in different conditions, the accuracy indicators of the upscaling coefficients throughout these 368 in situ sites were summarized as the histograms (Fig. 3).



300

Figure 3. The scatter plots between the upscaling results with the upscaling models and the coarse pixel
 scale reference. Only parts of the results are shown for conciseness. Specifically, only one in situ site is
 shown for each land cover type.

Based on the results presented in Fig. 4, it can be seen that the overall accuracy of 304 the upscaling coefficients is satisfactory. The biases are concentrated in 0, and more 305 than 90% of them are within the range of ± 0.02 (Fig. 4(b)). The highest density of R² 306 is between 0.9 and 1 as shown in Fig. 4(c), and only a small part of the sites show a 307 relatively small R² of lower than 0.8 but larger than 0.5. Nevertheless, it should be noted 308 that those sites exhibit a more scattered distribution of RMSE values, with a maximum 309 of 0.1 and a minimum of 0.01 (Fig. 4(a)). The highest density is between 0.03 and 0.05 310 for RMSE. 311

329







312 Dids R²
313 Figure 4. Distribution of RMSE (a), Bias (b), and R² (c) of the upscaling coefficients. The histograms
314 presented here combine the results of the 368 in situ sites.

315 Given the fact that the accuracy of the upscaling models shows great variability, it 316 is necessary to explore the influence factors on the performance of the upscaling models. In this study, the effect of land cover type, elevation, and spatial heterogeneity were 317 considered. The influence of spatial heterogeneity on the accuracy of the upscaling 318 model is displayed in Fig. 5. The RMSE exhibits a significant positive correlation with 319 spatial heterogeneity (Fig. 5(a)), with superior performance often observed in areas with 320 lower spatial heterogeneity. Similarly, the R² of different sites typically decreases with 321 the increase of spatial heterogeneity (Fig. 5(b)). It is worth noting that when the spatial 322 heterogeneity exceeds 0.1, the model's stability fluctuates considerably, indicated by 323 the larger height of the boxplots of RMSE and R². Based on these results, it can be seen 324 that spatial heterogeneity has enormous implications for the performance of the 325 upscaling models. One possible reason is that the assumption of a linear relationship 326 between the subpixel albedo of other locations and the subpixel albedo containing the 327 in situ site cannot be satisfied over the surface with large spatial heterogeneity. 328





331 heterogeneity. Three different degrees of spatial heterogeneity are marked by different colors.





The variation of RMSE and R^2 of upscaling models with elevation is shown in Fig. 332 6. It is interesting to see that the upscaling model presents the highest accuracy for 333 elevations below sea level, with the lowest RMSE around 0.3 and the highest R^2 334 exceeding 0.95. By contrast, the model's performance was the worst for elevations 335 above 2500 meters, displaying the highest RMSE and the lowest R². However, for the 336 337 areas with elevation between 0 m and 2500 m, the variation trend of RMSE and R^2 with elevation is not obvious. For instance, a slightly decreasing trend can be found as 338 elevation increase from 0-200 m to 500-1500 m. But a slight increasing trend appears 339 as elevation increase from 500-1500 m to the area above 2500 m. Except for the areas 340 below 0 m, both the RMSE and R² show a wide fluctuation range, indicated by the large 341 342 height of their boxplots. These results demonstrate that the accuracy of the upscaling model is controlled by many other factors. The good performance of the upscaling 343 model over the area below sea level may be attributed to the small spatial heterogeneity 344 as shown in Fig. 7. It can be found that the area below sea level is featured by small 345 spatial heterogeneity less than 0.1. 346



347 348

Figure 6. The variations of RMSE (a) and R^2 (b) of with elevation.

As illustrated in Fig. 7, it is evident that the RMSE exhibits a noticeable upward 349 trend with increasing spatial heterogeneity at each elevation level. This suggests that 350 spatial heterogeneity plays a dominant role in determining the performance of the 351 upscaling models.. Nevertheless, the influence of spatial heterogeneity seems to be 352 related to the elevation. From Fig. 7(a), it can be seen that the difference in RMSE 353 between different spatial heterogeneity levels tends to be larger with the increase of 354 elevation. A similar phenomenon can be found in R². For each specific level of spatial 355 heterogeneity, the variation trend of RMSE and R² with elevation is not the same. For 356 areas with small spatial heterogeneity (<0.1), their variation trend is not significant. By 357 contrast. For the area with strong spatial heterogeneity (>0.3), the increasing/decreasing 358 trend of RMSE/R² with elevation is obvious, especially for the area above 200 m. 359







362 Figure 7. The plots show combined results of the RMSE (a) and R^2 (b) variations based on elevation and



363 spatial heterogeneity.

364

Figure 8. The variations in RMSE (a) and R² (b) are dependent on landcover, different colors refer to the
 fifteen different land cover types.

The influence of land cover types on the accuracy of the upscaling model is displayed in Fig. 8. It shows that the model's accuracy is relatively poor on EBF, with most of the RMSEs exceeding 0.05 and R^2 below 0.90. By contrast, the model performs best on bare land (BSV), with the lowest RMSE close to 0.02 and a relatively higher





 R^2 around 0.97. The accuracy of the model over the remaining surface cover types was 371 similar, with RMSE and R² values around 0.05 and 0.90, respectively. The boxplots of 372 RMSE at URB, CVM, and BSV present a narrow range of values, indicating a relatively 373 stable model performance. Conversely, the boxplots of the RMSE at ENF and EBF 374 375 show a large range of values, indicating significant variability in model performance in 376 these areas. In fact, the influence of land cover types is also associated with the effect 377 of spatial heterogeneity. As shown in Fig. 9, the BSV is featured by small spatial heterogeneity. By contrast, a considerable number of sites over EBF show a strong 378 spatial heterogeneity. 379



380

Figure 9. The plots show combined results of the RMSE (a) and R² (b) variations based on land cover
and spatial heterogeneity.

As shown in Fig. 9(a), the RMSE of the upscaling model basically presents an increasing trend with spatial heterogeneity over each land cover type, further indicating the dominant role of spatial heterogeneity in determining the accuracy of the upscaling model. The outliers at OSH, URB, CVM, and WAT are attributed to the small number of sites with relatively larger spatial heterogeneity. Nevertheless, the influence of spatial heterogeneity show dependence on land cover type, which is most significant on GRA.

389 4.2 The accuracy of the pixel scale ground truth

390 Since there are a considerable number of in situ sites, the accuracy of pixel scale





ground truth was summarized as the boxplots (Fig. 10). For comparison purposes, the 391 errors of single site measurements when they were directly used as the pixel scale 392 reference were also calculated and summarized as the boxplots. It can be seen that the 393 errors of pixel scale ground truth show a slight variation with spatial heterogeneity, with 394 the median RRMSE ranging from 29.44 % to 31.51% and then to 29.35 %, resulting in 395 the overall RRMSE of 29.89 %. It is important to note that this variation pattern is not 396 397 the same as the accuracy of the upscaling model, which shows a monotonous decrease trend with the increase of spatial heterogeneity. The wide range of the boxplots shows 398 that the accuracy of the pixel scale ground truth is also influenced by other factors. 399

Indeed, the errors of the pixel scale ground truth are not negligibly small. 400 Nevertheless, its accuracy is consistently better than the single site measurements over 401 the surfaces with different levels of spatial heterogeneity as shown in Fig. 10. The 402 smaller RRMSE of pixel scale ground truth (29.89 % vs. 33.34 %) indicates that this 403 dataset can improve the accuracy of pixel scale reference data a lot compared to the 404 single site measurements. Nevertheless, the degree of the improvement depends on the 405 406 situation, which is the most significant over the sites with the strongest spatial heterogeneity, with the RRMSE decreasing from 36.62 % to 29.35 %. The in situ sites 407 with medium spatial heterogeneity follow, with the RRMSE decreasing from 35.12 % 408 to 31.51 %. The improvements are the smallest over the sites with the smallest spatial 409 heterogeneity, with the RRMSE decreasing from 32.69 % to 29.44 %. Hence, it can be 410 411 concluded that the degree of improvements of this dataset shows an increasing trend with spatial heterogeneity. Furthermore, the accuracy of the pixel scale ground "truth" 412 dataset is more stable than that of the single site measurements, indicated by the smaller 413 height of the boxplots of the former. 414







415

Figure 10. The boxplots of RRMSE of pixel scale ground truth and single site measurements. The
boxplots are categorized by different degrees of spatial heterogeneity and overall accuracy. The median
of the boxplots is indicated by the numbers around the plots.

The aforementioned results confirm the efficacy of our pixel scale ground "truth" dataset in different scenarios, which is superior to single site measurements, whether for the sites with higher or lower spatial heterogeneity. The improvement of this dataset is more significant over the heterogeneous sites. Hence, it is highly helpful over heterogeneous surfaces in the validation or bias correction of satellite albedo products.

424 **4.3** The usage of the pixel scale ground "truth" dataset

Validation is important for both the satellite product producers and end users as it 425 provides a quantitative evaluation of the advantage and disadvantages of satellite 426 products. However, in situ measurements, locations, coverage, scaling, and 427 representation, are rather diverse, resulting in a different impression of accuracy. 428 Consequently, most of the validation results are not directly comparable, which further 429 limits the potential utility of satellite products. As pointed out by GCOS, one possibility 430 for addressing this challenge is through a unified and systematic validation based on a 431 unified, standard, and consistent reference dataset. The pixel scale ground truth was 432 obtained through a standardized operational procedure based on a large number of in 433 situ sites' measurements distributed across the globe. Such standardization enables the 434 fair comparison between the accuracy of different satellite products of the same ECV, 435 thus providing a foundation for the coordinated utilization of diverse satellite albedo 436 products and maximizing their potential capabilities. Fig. 11 presents an example for 437







438 the validation of MCD43A3 V0061 based on this pixel scale ground dataset.

Figure 11. The scatter plots between the MCD43A3 and pixel scale ground truth (green dots) as well asthe scatterplots between CDF-corrected MCD43A3 and pixel scale ground truth (brown dots).

Apart from validation, the pixel scale ground "truth" dataset can also be used as a 442 reference to correct the bias of satellite albedo products. There have been many kinds 443 of bias correction models (Wang et al., 2022). In this study, the CDF (cumulative 444 distribution function) method (Calheiros et al., 1987) was utilized to correct the bias in 445 MCD43A3 V0061 for example purposes. As shown in Fig. 10, the bias correction 446 generally improves the accuracy of satellite albedo products, and the improvement is 447 particularly significant for regions with strong heterogeneity (Fig. 10(d-f)). Hence, it is 448 reasonable to consider that this dataset has the potential to further improve the quality 449 of satellite albedo products, particularly in regions with strong surface heterogeneity. 450

451 **5. Data availability**

The processed coarse pixel scale ground "truth" dataset is openly accessible and can be obtained from Zenodo (https://doi.org/10.5281/zenodo.8008455; Pan et al., 2023). The dataset files are available in the machine-readable data format (.tab) and have been conveniently organized into separate folders for Ameriflux, Euroflux, BSRN, SURFRAD, HiwaterWSN, and Huailai station, facilitating easy accessibility and



457 utilization.

458 6. Conclusion

Both validation and correction of remote-sensing albedo products are essential to 459 promote the rational utilization of satellite albedo products in various scientific 460 applications. Typically, in situ single site-based albedo measurements with a 461 widespread are the backbone of validation and correction. However, since satellite-462 based albedo and tower-based albedo are generated at different spatial scales, direct 463 464 comparison is only valid over specific homogeneous land surfaces. Nevertheless, spatial heterogeneity is a basic feature of most land surfaces, resulting in the limited 465 spatial representativeness of single site measurements. Thus, the most critical aspect of 466 467 validation/ correction is obtaining the pixel scale ground truth albedo based on in situ site measurements. 468

However, the ways of obtaining pixel scale ground truth are rather diverse 469 regarding the in situ measurements, locations, coverage, scaling, and representation, 470 resulting in different accuracy of pixel scale "truth" dataset. Consequently, most of the 471 472 validation/correction results are not directly comparable, which further limits the further use of satellite products. A suite of continuous, independent, and representative 473 pixel scale ground "truth" dataset of surface albedo is needed. To the best of our 474 knowledge, there is still a lack of such a kind of dataset with global coverage. To fill 475 this gap, we have developed a coarse pixel-scale ground truth albedo dataset based on 476 ground measurements of 368 in situ sites from the sparsely distributed observation 477 networks (e.g., SURFRAD, BSRN, and Fluxnet) and an upscaling model specified for 478 479 the single site measurements.

480 The applicability of the upscaling model to the in situ site measurements was first comprehensively on a global scale. The overall accuracy of the upscaling coefficients 481 is satisfactory, with the biases concentrated around 0. 90 % of these biases are within 482 the range of ± 0.02 . The influence factors on the performance of the upscaling models 483 were also explored. Spatial heterogeneity plays a dominant role in determining the 484 performance of the upscaling models. But the influence of spatial heterogeneity is 485 related to the elevation and land cover type, which tends to be larger with the increase 486 of elevation and over the land cover type of GRA. 487

The accuracy of the pixel scale ground "truth" dataset was also comprehensively evaluated. The errors of pixel scale ground truth show a slight variation with spatial heterogeneity, with the median RRMSE ranging from 29.44 % to 31.51 % and then to





29.35 % over the surfaces with small, medium, and large spatial heterogeneity, 491 respectively, resulting in the overall RRMSE of 29.89 %. Although the errors of the 492 pixel scale ground truth are not negligibly small, its accuracy is consistently better than 493 the single site measurements over the surfaces with different levels of spatial 494 heterogeneity. The smaller RRMSE of pixel scale ground truth compared to the single 495 site measurements (29.89 % vs. 33.34 %) indicate that this dataset can improve the 496 accuracy of pixel scale reference data a lot. The degree of improvements of this dataset 497 shows an increasing trend with the spatial heterogeneity. Furthermore, the accuracy of 498 the pixel scale ground "truth" dataset is more stable than that of the single site 499 measurements. 500

The pixel scale ground truth was obtained through a standardized operational procedure. Such standardization enables the fair comparison between the accuracy of different satellite products of the same ECV. This newly introduced dataset serves as a remedy to the inadequacy and inconsistency of reference data currently employed in validation/correction efforts, thereby paving the way for the coordinated use of various satellite albedo products and unlocking the full capacity of different albedo products.

507 Author contributions

FP, XW was responsible for the main research ideas and writing the manuscript. JW,
DY, QX contributed to the data collection. XW, RT, QZ, ZL, JW have supported the
work with formal analysis, XW contributed to the manuscript organization. All authors
have worked on the writing, particularly for review and editing.

512 **Competing interests**

513 The contact author has declared that none of the authors has any competing interests.

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