



Microwave radiometry experiment for snow in Altay
 China: in situ time series of data for electromagnetic and

3 physical features of snow pack and environment

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19 Abstract. Snow depth is a key parameter in climatic and hydrological systems. Passive microwave 20 remote sensing, snow process model and data assimilation are the main methods to estimate snow depth 21 in large scale. The estimation accuracies strongly depend on input of snow parameters or characteristics. 22 Because the evolving processes of snow parameters vary spatiotemporally, and are difficult to accurately 23 simulate or observe, large uncertainties and inconsistence exist among existing snow depth products. 24 Therefore, a comprehensive experiment is needed to understand the involution processes of snow 25 characteristics and their influence on microwave radiation of snowpack, evaluating and improving the 26 snow depth and SWE retrieval and simulation methods. An Integrated Microwave Radiometry Campaign 27 for snow (IMCS) was conducted at the Altay National Reference Meteorological station (ANRMS) in 28 Xinjiang, China, during snow season of 2015/2016. The campaign hosted a dual polarized microwave 29 radiometer operating at L, K and Ka bands to provide minutely passive microwave observations of snow 30 cover at a fixed site, daily manual snow pit measurements, ten-minute automatic 4-component radiation 31 and layering snow temperatures, covering a whole snow season of 2015/2016. The measurements of 32 meteorological and underlying soil parameters were requested from the ANRMS. This study provides a 33 summary of the obtained data, detailing measurement protocols for microwave radiometry, in situ sow 34 pit and station observation data. A brief analysis of the microwave signatures against snow parameters is 35 presented. A consolidated dataset of observations, comprising the ground passive microwave brightness 36 temperatures, in situ snow characteristics, 4-component radiation and weather parameters, was achieved 37 in the National Tibetan Plateau Data Center, China. The dataset is unique in providing continuously daily 38 snow pits data over a snow season at a fixed site and matched microwave brightness temperatures, 39 radiation and meteorological data. The dataset is expected to serve for the evaluation and development 40 of the microwave and optical radiative transfer models and snow evolution process models. The

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41 consolidated data are available at https://data.tpdc.ac.cn/en/ (doi: 10.11888/Snow.tpdc.270886) (Dai,

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44 Key words: Snow, Microwave radiometry, Snow pit, Experiment

# 45 1 Introduction

46 Seasonal snow cover plays a critical role in climate and hydrological systems (Cohen, 1994; Ding 47 et al., 2020; Barnett et al., 2005; Immerzeel et al., 2010) by its high albedo, thermal insulation, fresh 48 water reserves and its phase change process. Snow cover can be accurately identified by optical remote 49 sensing. However, the snow surface albedo was controlled by snow characteristics (Aoki et al., 2003 and 50 2000), and the variations in snow characteristics cause the uncertainties of albedo estimation. Snow depth 51 and snow water equivalent (SWE) are currently estimated using passive microwave in global and 52 regional scales (Pullianen et al., 2020; Tedesco and Narvekar, 2010; Jiang et al., 2014; Che et al., 2008). 53 Although several global and regional snow depth and SWE products have been released, large 54 uncertainties exist in these products because of the spatio-temporal variations in snow characteristics 55 (Xiao et al., 2020; Mortimer et al., 2020; Che et al., 2016; Dai et al., 2012). Therefore, the observation 56 on electromagnetic and physical parameters of snowpack are necessary to improve understanding the 57 electromagnetic radiation process of snowpack to enhance the estimation accuracy of snow surface 58 albedo and snow depth.

59 To evaluate and improve snow depth and SWE retrieval methods from passive microwave remote 60 sensing observations and to combine remote sensing technologies with modeling and data assimilation 61 methods for producing the most accurate products, a few large or systematic experiments or campaigns 62 had been conducted on electromagnetic and physical characteristics measurement of snow cover. The 63 Cold Land Processes Field Experiment (CLPX) (https://nsidc.org/data/clpx/index.html) was one of the 64 most well-known experiments, which was carried out from winter of 2002 to spring of 2003 in Colorado, 65 USA (Cline et al., 2003). During the campaign, snow pits were collected at the February and March of 66 2002 and 2003 to coincide with the airborne and ground remote sensing observations. NASA SnowEx 67 campaign (https://nsidc.org/data/snowex) was conducted in 2017 in Colarado to develop/test algorithms 68 for measurement of SWE in forested and non-forested areas by providing multi-sensor observations of 69 seasonally snow-covered landscapes (Brucker et al., 2017). The campaign is still ongoing and will be 70 conducted in other areas with different snow conditions. In northern Canadian region, mobile sled-71 mounted microwave radiometers were deployed in forest, open and lake environments from November 72 2009 to April 2010 and snow characteristics within the footprints of radiometers were measured to 73 improve understanding the influence of snow characteristics on brightness temperatures (Derksen et al., 74 2012; Roy et al., 2013). These microwave experiments were of mobile observation. In these experiments, 75 there were multiple observation sites for different land cover, but relative short temporal range. The snow 76 pit observations could be used for evaluating snow microwave emission model in different land cover 77 (Tedesco and Kim, 2006; Royer et al., 2017), but they did not exhibit the involution of snow parameters. 78 In the Arctic region, the Nordic Snow Radar Experiment (NoSREx) campaign was conducted at a 79 fixed field in Sodankylä, Finland, during 2009 ~ 2013 (Lemmetyinen et al., 2016). This experiment 80 provided a continuous time series of active and passive microwave observations of snow cover at a 81 representative location of the Arctic boreal forest area covering a whole winter season and matched snow 82 pit observations were made weekly. In Asia, snow pit work of 3 or 4-day intervals was conducted





83 simultaneously with radiation budget observations during winter of 1999/2000 and 2000/2001 to analyze 84 the effects of snow physical parameters on albedo (Aoki et al., 2003). The NoSREx and Japan radiation 85 experiments were of fixed field observation, which provided longer time series of data. These 86 experiments were conducted in deep snow area, and the week-interval observation could reflect the 87 general evolution process of snow characteristics, but might miss some details. Furthermore, in the area 88 with snow cover duration within 4 months, the week-interval observation hardly depicts the change 89 details.

90 To comprehensively understand the evolutions of snow characteristics and their influence on passive 91 microwave brightness temperatures and radiation budget, an integrated experiment on snow was 92 conducted during a whole snow season, in Altay, China. The experiment was designed to cover periods 93 from snow-free conditions to eventual snow melt-off during 2015/2016. The microwave radiometry 94 measurements at L, K and Ka bands for multiple angles were complemented by a dual-polarized 95 microwave radiometer with 4-component radiation and daily in situ observations of snow, soil and 96 atmospheric properties, using both manual and automated methods. The data of electromagnetic and 97 physical parameters were further consolidated and organized to be easily read and utilized.

98 The dataset is unique in providing continuously daily snow pits data over a snow season at a fixed 99 site and matched microwave brightness temperatures, radiation and meteorological data. In the next 100 section, the experiment location, parameters, parameter measurement protocols are described; section 3 101 introduces the consolidated data which was released on the National Tibetan Plateau Data Center, China; 102 section 4 presents content of brightness temperature, 4-component radiation, snow pit data, soil 103 temperature and moisture, and meteorological data; section 5 discusses the possible application; and 104 finally the conclusions are summarized in section 6.

# 105 2 Description of experiment setup

# 106 2.1 Measurement location

107 The Integrated Microwave Radiometry Campaign for snow (IMCS) was performed during snow season of 2015/2016 (from November 27, 2015 to March 25, 2016) at the Altay National Reference 108 Meteorological station (ANRMS) (N47°44'26.58", E 88°4'21.55") which is at the foot of Altay 109 110 mountain in the northwest China, and approximate 6 km distance from the mountain (Figure 1). Altay 111 mountain, running northwest and southeast, is at the junction of China, Russia, Mongolia and Kazakhstan, 112 and the elevation is up to 3000 m, providing snow water resources for the four countries. The average 113 annual maximum snow depth measured in this station is approximately 40 cm, and the maximum value 114 is over 70 cm. In the southwest of Altay mountain, crop land and desert with flat terrain are the dominant 115 land covers. Snow cover is critical fresh water for the irrigation in this area. In this experiment, measurements included microwave radiometry, 4-component radiation, snow pit and soil parameters. 116 117 The test site of this experiment was four neighboring bare rectangle fields in the ANRMS with their area 118 of 2500m<sup>2</sup> (black rectangle filed in Figure 1), 2500m<sup>2</sup> (pink rectangle field in Figure 1), 200m<sup>2</sup> (red rectangle field in Figure 1) and 400 m<sup>2</sup> (blue rectangle field in Figure 1), respectively. 119

120 In the pink field, the ground-based microwave radiometer was set up in the middle place of this 121 field, facing south to collect brightness temperatures of snow cover. The black field behind the 122 microwave radiometers (north of the radiometers) was for snow pit data collection. The microwave



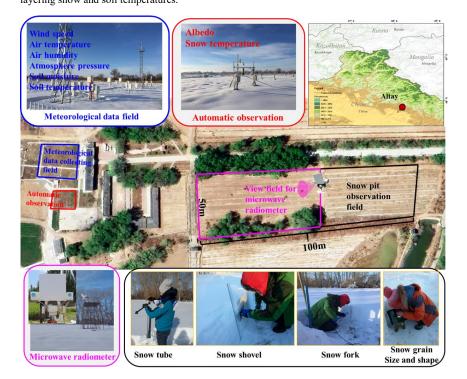


radiometer observations and snow pit data collection were conducted by Northwest Institute of Eco Environment and Resources, Chinese Academy of Science (NIEER) from November 27, 2015 to March
 25, 2016 (After March 25, 2016, snow melted out).

The blue field is for meteorological measurements including wind speed, wind direction, air temperature, air wetness, air pressure, precipitation, layering soil temperature, soil moisture among others. These parameters were automatically obtained from instruments, and the instruments setup and data collection were operated by ANRMS. This station also has daily manual observation in snow depth and SWE. In this experiment, we requested the wind, air pressure, air wetness, soil temperature and moisture data during this experiment from ANRMS.

132The red field was designed for automatic measurement of layering snow temperatures, snow density,133SWE, snow depth, and albedo. These automatic measurement instruments were installed and maintained134by NIEER, and started working from 2013. However, during the experiment, the instruments for snow135density and SWE did not work, and we only collected layering snow temperatures and 4-component136radiation from November 27, 2015 to March 25, 2016.

Because the four observation fields are located within the domain of the station and the distance between each other are less than 100m, the snow characteristics and soil and weather conditions are thought to be the same. Ultimately, the collected data in this study include ground-based brightness temperatures, 4-component radiation, snow pit data, meteorological data and automatically observed layering snow and soil temperatures.



143Figure 1: Location of the Altay National Reference Meteorological station (ANRMS) in the Asia and the144distribution of three experiment fields in the ANRMS. The black rectangle represents the field used for snow145characteristics (approximately 40 m × 50 m) including layering thickness, snow density, snow grain size and





146 shape, and microwave radiometers (approximately 60 m × 50 m) observations. The blue rectangle is the field 147 for Meteorological and soil data collection operated by the ANRMS. The red rectangle field is for the 148 automatically observation of the snow temperature, SWE, 4-component radiation and snow depth, designed 149 by Northwest Institute of Eco-Environment and Resources, Chinese Academy of Science (NIEER). Note: The 150 map in the up right corner is ArcGIS self-contained map.

# 151 2.2 Measurement parameters and instruments

The brightness temperatures at 1.4, 18.6, 36.5 GHz for both polarization (Tb1h, Tb1v, Tb18h, Tb18v, 152 153 Tb36h, Tb36v) were automatically collected using a six-channel dual polarized microwave radiometer 154 **RPG-6CH-DP** (Radiometer Physics GmbH. Germany, 155 https://www.radiometerphysics.de/products/microwave-remote-sensing-instruments/radiometers/). The technical specifications of the RPG-6CH-DP are described in Table 1. The RPG-6CH-DP contains a 156 157 built-in temperature sensor which can measure air temperature. The automatically data collection 158 frequency was set as 1 minute.

159

# 160 Table 1. Technical Specifications of the RPG-6CH -DP Microwave Radiometer.

Parameter	Value			
Manufacturer	Radiometer Physics GmbH			
System noise temperatures	<900 K			
Bandwidth	400MHz (20MHz for 1.4 GHz)			
System stability	0.5 K			
Dynamic range	0~400 K			
Frequencies (GHz)	1.4, 18.7, 36.5			
Polarizations	V, H			
Internal calibration	Internal Dicke switch and software control for			
Internal calibration	automatic sky tilt calibration			
Receiver and antenna thermal stabilization	< 0.015 K			
Antenna sidelobe level	< -30 dBc			
Optical resolution (HPBW)	6.1° (11° for 1.4 GHz)			
Incidence angle	0~90°			
Azimuth angle	360°			
ç	* , *			

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162 The microwave signatures from snowpack vary with snow characteristics, soil and weather 163 conditions. The snow characteristics were obtained by manual snow pit measurements in the black field, 164 including layering snow thickness, layering grain size, snow density, and snow temperatures. These data 165 were daily collected during 8:00-10:00 am local time, from November 27, 2015 to March 25, 2016, 166 except 7 days (please see Table 1). Although the snow temperatures were manually measured at snow pits, the automatically collected snow temperatures in the red field were utilized in this study, because 167 168 the temperature measured at snow pits could not reflect the natural temperature profile when the snow 169 pits exposed to the air. In the red field, the 4-component radiation was automatically measured by NR01 170 manufactured by Hukseflux, and layering snow temperatures was measured by Campbell 109S temperature sensors, respectively. The soil and weather parameters are of routine observations at 171





- 172 ANRMS, and were obtained through request from ANRMS. The instruments used for soil and weather
- 173 parameters observations are produced by China Huayun Meteorological Technology Group corporation.

174 The measurement parameters and their measurement instruments are listed in table 2.

176	Table 2. Variables collected in MRESC and the observation equipment, observation time and frequencies.
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Parameter	Equipment/Method	Layering style	Observation time or frequency	Absent date
Tb1h, Tb1v, Tb18h, Tb18v, Tb36h, Tb36v and air temperature	Microwave radiometers RPG- 6CH-DP/ Automatically	5 meters above ground	minutely	no
Layer thickness (cm)	Ruler / Manually	Natural layering	local time 8:00- 10:00 am	no
Snow density (g/cm3)	Snow tube / Manually	Whole snowpack	local time 8:00- 10:01 am	no
Snow density (g/cm3)	Snow shovel/ Manually	Every 10 cm	local time 8:00- 10:02 am	1/2 2/2017
Snow density (g/cm3) and liquid water content (%)	Snow fork/ Manually	At interval of 5 cm from the 5-cm height	local time 8:00- 10:03 am	1/2-3/2016, 2/20/2016
Snow grain size (mm)	Anyty V500IR/UV/ Manually	Natural layering	local time 8:00- 10:04 am	12/24/2015, 12/31/2015, 1/
Snow grain shape	Visually identification/ Manually	Natural layering	local time 8:00- 10:05 am	1-3/2016, 1/23/ 2016, 2/20/2016
Snow temperature(°C)	Temperature sensors (Campbell 1098) / Automatically	0 cm, 5 cm, 10 cm, 15 cm, 25 cm, 35 cm, 45 cm, and 55 cm	Ten-minute	no
4-component radiation (W/m2)	Component Net Radiometer (NR01) / Automatically	6 feets above ground	Ten-minute	no
Soil temperature(°C)	Soil temperature sensor (China Huayun) / Automatically	5cm below soil/snow surface (-5cm)	Hourly	no
Soil moisture (%)	Soil moisture sensor (DZN3, China Huayun) / Automatically	10cm below soil/snow surface (-10cm)	Hourly	no
Air temperature(°C)	Thermometer screen (China Huayun) / Automatically	6 feets above ground	Hourly	no





Air pressure (Pa)	Thermometer screen (China Huayun) / Automatically	6 feets above ground	Hourly	no
Air humidity (%)	Thermometer screen(China Huayun) / Automatically	6 feets above ground	Hourly	no
Wind speed (m/s) and direction	Wind sensor(China Huayun) / Automatically	10 m above ground	Hourly	no

# 177

### 178 **2.3 Measurement methods**

179 In this experiment, the measurements include microwave radiometry observation to collect 180 brightness temperature, snow pit observation to collect snow physical parameters, automatic observation 181 to collect 4-component radiation and snow temperatures, and meteorological observation which contains 182 weather data and soil data.

183 **2.3.1. Microwave radiometry** 

Before snow season, a platform with height of 5 m, length of 4 m and width of 2 m was constructed in the experiment field. A 4-m orbit was fixed on the platform. The RPG-6CH-DP was set up on the orbit and could be moved along the orbit. This radiometer was sky tipping calibrated. In the clear sky conditions, the sky brightness temperatures are approximately  $7.8\pm1$ K and  $15.7\pm0.7$ K at 1.4 GHz for horizontal polarization and vertical polarization, respectively; those were approximately  $29.7\pm0.3$  K and  $29.3\pm0.9$  K at 18.7GHz and 36.5 GHz, respectively.

190 Generally, the radiometers were fixed in the middle place of the orbit to observe snow cover with 191 incidence angle of 45°. Multi-angle observations were conducted on seventeen days (Dec 3, Dec 19, Dec 192 30, Jan 3, Jan 8, Jan 13, Jan 18, Jan 23, Jan 28, Feb 3, Mar 3, Mar 10, Mar 15, Mar 22, Mar 26, Mar 28, 193 Mar 31) when the radiometer was set to scan the ground at different incidence angles at two ends of the 194 orbit and the middle place of the orbit. Although the view fields of the antennas for 1.4 GHz, 18 GHz 195 and 36 GHz did not completely overlap, the measured results showed that the brightness temperatures observed by radiometers at the left, middle and right of the orbit varied less than 1 K. Therefore, the 196 197 snow and soil characteristics presented homogeneous distribution within the view files of the three 198 antennas.

199

# 200 2.3.2 Snow pit measurement

201 The physical snow parameters were measured at snow pits. These parameters included snow depth, 202 snow density, snow layering, layering grain size and shape type, layering snow density.

The first step of snow pit measurement is making snow pit. In the black field, snow pits were made every day. A spade was used to excavate snow pit. The length of the snow pit profile was approximately 205 2m to make sure all parameters were measured from unbroken snowpack. The width of the snow pit was 206 approximately 1m for observers to conveniently observe. The snow pit section was made as flat as 207 possible using a flat shovel or ruler. When the snow profile expose to the air for a long time, the snow 208 characteristics will be influenced by environment and will be different from the natural snow 209 characteristics. In order to make sure every observation conducted on natural snow pit, the snow pit was





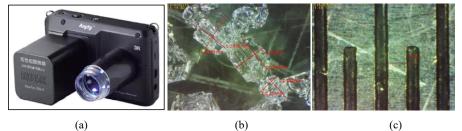
backfilled with the shoveled snow after finishing all observations, and the new snow pit in the following
 day was made at least 1-m distance from the last snow pit.

212 The second step of snow pit measurement was recording the natural stratification. After finishing a

snow pit, the natural stratification of snowpack was firstly visually determined against a ruler, and the snow thickness of every layer was recorded.

215 The third step was measuring grain size and shape type in each layer, whilst snow density was measured at constant interval. The grain size and type within each natural layer were estimated visually 216 217 from a microscope with a camera named "Anyty V500IR/UV" (Figure 2a). A software matched the 218 microscope was used to measure the grain size. The grain type was determined based on Fierz et al. 219 (2009). In this experiment, we utilized the length of longest axes and the length of shortest axes to 220 describe grain size (Figure 2b). When using the software to measure the grain size, a reference must be 221 needed. In this experiment, the minimum scale "0.5mm" of a ruler was used as reference (Figure 2c). We 222 adjusted the focus of the camera to make sure the grains at the clearest status in camera to take photos, 223 and the photo of ruler scale was taken at the same focus. If the thickness of one layer was less than 10 224 cm, measurements were performed at the top and bottom of the layer. If the thickness was greater than 225 10 cm, measurements were performed at the top, middle, and bottom of the layer. For each layer, at least 226 5 photos were taken, and at least 10 typical grains were chosen to measure the longest axes length and the shortest axes length in the photos of each layer. Then, in each layer, there were at least 10 groups of 227 228 the longest and shortest axes length were obtained, and the final grain size was the average values. Figure 229 Al presents an example of the original photos of grains in each layer, and Table Al shows the matched 230 record of longest and shortest axis length.

231



232 233 234

Figure 2: Picture of microscope "Anyty V500IR/UV" (a), the measured longest axes lengths and shorteast axes length of particles (b), and the reference ruler scale (c).

235 236

237 Snow density was measured using three instruments: snow tube, snow shovel and Snow Fork 238 (Figure 3). The snow tube instrument, designed by Chinese Meteorological administration, contains a 239 metal tube with the base area of 100 cm<sup>2</sup> and the length of 60 cm, and a balance (figure 3a). It was utilized 240 to measure the snow density of a whole snowpack by weighing the snow sample. The snow shovel is a 241 1500 cm<sup>3</sup> wedge-type sampler, and its length, width and height are 20 cm, 15 cm, and 10 cm, respectively 242 (figure 3b). It was utilized to measure layering snow density every 10 cm (0-10 cm, 10-20 cm, 20-30 243 cm...). The Snow Fork is a microwave resonator that measures the complex dielectric constant of snow, 244 and adopts a semi-empirical equation to estimate snow density and liquid water content based on the 245 complex dielectric. The Snow Fork (figure 3c) was utilized to measure snow density and liquid water 246 content at 5-cm intervals from the 5-cm height over the snow/soil interface (5cm, 10cm, 15 cm, 20cm...). 247 Table A2 is an example of record table for snow density. Three times of observations were conducted for





248 249	every layer.		
250			
251	(a)	(b)	(c)
252	Figure 3: Three instruments	for snow density: Snow tube (a), Snow shovel (b), a	
253			
254	2.3.2 Automatic radiation	and temperature measurement	
255	Layering snow temper	atures were collected using Temperature sensor	s at the red field instead of
256	manual observation from si	now pits. The temperature sensors had been set	up on a vertical pole which
257	was vertically inserted in th	he soil (Figure 4). The heights of the sensors are	0 cm, 5 cm, 10 cm, 15 cm,
258	25 cm, 35 cm, 45 cm, and 5	5 cm. The snow temperatures at these heights we	ere collected every minute.
259	TheNR01 net radiome	ter was set up to measure the energy balance bet	tween incoming short-wave
260	and long-wave Far Infrare	ed radiation versus surface-reflected short-wave	and outgoing long-wave
261	radiation. The range of sho	rt wave is 285~3000nm, and the range of long	wave is 4.5~40um. The 4-
262	component radiation was a	atomatically recorded every ten minutes. In addi	tion, the sensor is equipped
263	with a Pt100 temperature se	ensor for parallel recording of the sensor temperat	ture.







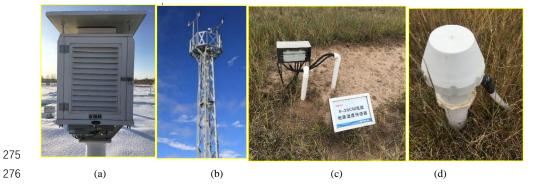
- 264
- 265 Figure 4: Set up of temperature sensors and CNR4 in the red field.
- 266

# 267 2.3.3 Meteorological observation

268The meteorological data requested from the ANRMS include air temperature, air pressure and269humidity, wind speed, soil temperature at -5cm, -10 cm, -15cm and -20 cm and soil moisture at -10 cm270and -20 cm.

271The air temperature, pressure and humidity were collected using temperature and wetness sensor in272thermometer screen, the wind speed and direction were measured using wind sensor set up at 10 m on a

- 273 tower. Soil moisture and temperature were automatically measured using moisture sensor and
- 274 temperature sensor. Figure 5 depicts the instruments for these observations.



- 277 Figure 5: Instruments for observation of air temperature and wetness (a), wind speed (b), soil temperature
- 278 (c) and soil moisture (d).





# 279 3 Description of consolidated IMCS data

280	The microwave brightness temperature, snow parameters, meteorological data were recorded in
281	different formats, and the observation frequencies and times were different. These data must be
282	reorganized or consolidated for easily usage. The values from the three-time measurements for snow
283	density in each layer were averaged to obtain the final snow density. The length of the longest and shortest
284	axes of particles in each photo were measured using the software. The average lengths of longest and
285	shortest axes from all photos in each layer were obtained as the final grain size.
286	The ground-based brightness temperatures and layering snow temperatures were automatically
287	collected in uniform format. The weather and soil data requested from ANRMS had been consolidated.
288	Finally, the provided datasets are as following and table 3 describes the contents of each dataset.
289	1) Brightness temperatures data:
290	1 Minutely calibrated brightness temperature at 1.4 GHz, 18 GHz and 36 GHz for both polarizations at
291	incidence angle of 45°. This data include date, time, brightness temperatures at the three bands for both
292	polarizations.
293	2 Seventeen groups of calibrated brightness temperature at 1.4 GHz, 18 GHz and 36 GHz for both
294	polarizations at different incidence angles (30, 35, 40, 45, 50, 55, 60°). This data include date, incidence
295	angles, brightness temperatures at the three bands for both polarizations.
296	2) Snow pit data:
297	1 Daily grain size data include date, snow depth, layering snow thickness, average longest axis, average
298	shortest axis of each layer. The data were stored in an excel sheet.
299	2 Daily snow density data include date, layering snow density using snow fork (snow density at different
300	heights, such as SF_5cm, SF_10cm, SF_15cm), snow density using snow tube, layering snow density
301	using snow shovel (such as SS_0-10cm, SS _10-20cm, SS _20-30cm, SS _30-40cm). The data were
302	stored in an excel sheet.
303	3 Ten-minute snow temperatures data include date, time, temperatures at different heights (such as
304	ST_0cm, ST_5cm).
305	3) Snow surface radiation data
306	1 4-component radiation data include date, time, short-wave incident radiation, short-wave reflected
307	radiation, long-wave infrared incident radiation, long-wave infrared reflected radiation, and sensor
308	temperature.
309	4) Meteorological data:
310	1 Hourly weather data include date, hour, air temperature, wetness and humidity wind speed and direction.
311	2 Hourly Soil data include date, hour, soil temperature at 5 cm, 10 cm, 15 cm and 20 cm, and soil moisture
312	at 10 cm and 20 cm.
313	These data were stored in four folders, and different parameters were stored in different excel table or
314	ascii files. The organization of these consolidated data were described in table 3.
315	
316	Table 3 Description of consolidated data

 Folder
 Data
 Store format
 Content

 Brightness
 Minutely brightness
 Ascii file
 Date, time, Tb1h, Tb1v, Tb18h, Tb18v, Tb36h, Tb36v

 temperature
 temperature
 Date, time, Tb1h, Tb1v, Tb18h, Tb18v, Tb36h, Tb36v



	Multi-angle brightness temperatures	Ascii file	Date, time, incidence angle, Tb1h, Tb1v, Tb18h, Tb18v, Tb36h, Tb36v
Snow pit data	Grain size	Excel sheet	Date, snow depth, th1, Lg1, Sg1, th2, Lg2, Sg2,th3, Lg3, Sg3, th4, Lg4, Sg4,th5, Lg5, Sg5, th6, Lg6, Sg6
	Snow density	Excel sheet	Date, Stube, SS_0-10, SS_10-20, SS_20-30, SS_30-40, SS_40-50, SF_5, SF_10, SF_15, SF_20, SF_25.SF_30.SF_35.SF_40.SF_45.SF_50.
	Snow temperature	Excel sheet	Date, time, snow depth, ST_0cm, ST_5cm, ST_15cm, ST_25cm, ST_25cm, ST_45cm, ST_55cm
Radiation data	4-component radiation	Excel sheet	Date, time, SR_DOWN, SR_UP, LR_DOWN, LR_UP, T_Sensor
Meteorological data	Hourly weather data Hourly soil moisture and temperature	Excel sheet Excel sheet	Date, hour, Tair, Wair, Pair, Win Date, SM_10cm, SM_20cm, Tsoil_5cm, Tsoil_10cm, Tsoil_15 cm, Tsoil_20cm

317 Note: th: snow thickness, Lg: long axis, Sg: short axis;

318 Stube: snow density observed using snow tube, SS: snow density observed using snow shovel, SF: snow density

319 observed using snow fork; ST: snow temperature; SR\_DOWN: downward short-wave radiation, SR\_UP: upward

320 short-wave radiation, LR\_DOWN, downward long-wave radiation, LR\_UP: upward long-wave radiation, T\_sensor:

321 sensor temperature; Tair: air temperature, Wair: air wetness, Pair: air pressure, Win: wind speed.

322

# 323 4 Overview of collected data from IMCS

# 324 4.1 Snow characteristics

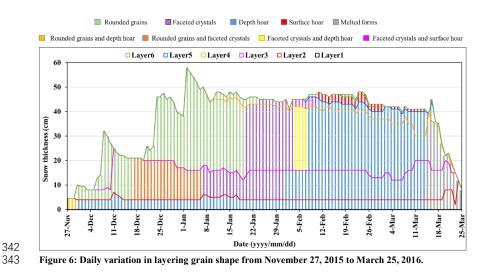
# 325 4.1.1 Layering grain size and grain shape

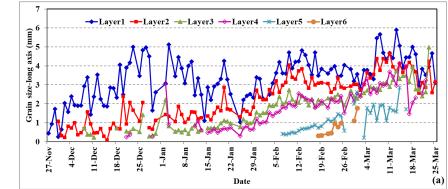
During winter of 2015/2016, snow cover began on 25 November of 2015, and ended on March 25 of 2016. During this snow cover duration, seven snowfall events occurred, and each snowfall formed one layer in snow cover on the ground, except the third event which presented a new layer on the second layer at the beginning, but the layering interface disappeared after several days and visually displayed as one layer (in gray in Figure 6). The fourth event was biggest of all, and the depth of snow cover exhibited decreasing with increase of snow density after the fourth event. Snow cover began melting on March 14 and snow depth declined to zero within 10 days.

333 Grain sizes within all layers exhibited increasing during the snow season, except in the bottom layer 334 where grain size experienced a decrease from December 28 to January 20 (Figure 7). In the vertical 335 profile, grain size increased from up to down with the snow age. The grain size of the fresh snow was 336 approximately 0.3 mm. The biggest long and short axis were up to 6 cm and 4 cm, respectively, which occurred within Layer 1 in the melting period. The length of short axes is approximately 0.7 of the length 337 338 of long axes. The grain shape generally developed from rounded grains to facet crystals, and then to 339 depth hoar. After March 13, 2016, the minimum air temperature increased to above 0, snowpacke accelerately melted, and the grain shape developed to melted forms. 340

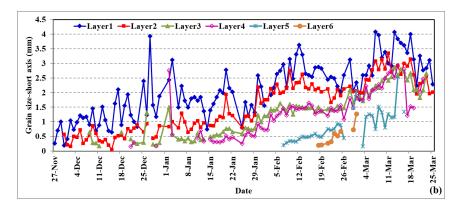






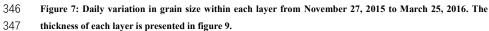


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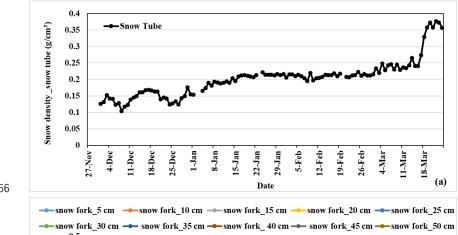
**4.1.2 Snow density** 

Snow densities measured by three different equipment shows that the density of fresh snow ranged

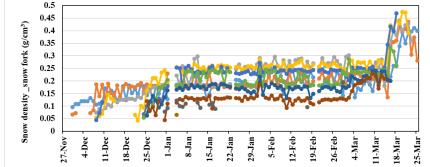




between 0.05~1.0 g/cm3 (Figure 8). The snow densities increased with snow age, and kept stable when the value arrived at 0.2~0.25g/cm3. From March 14 on, snow densities abruptly increased. The biggest value was beyond 0.45g/cm3. In the vertical profile, snow density increased from up to down in the accumulation phase, but in the stable phase, snow densities in the middle layers were larger than those in the bottom and upper layers. In the melting phase, snow densities in all layers showed little difference.

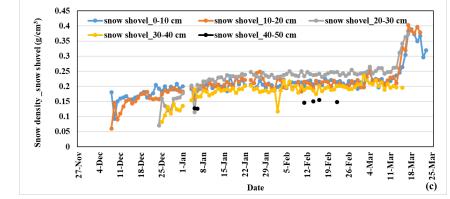






Date

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Figure 8: Daily variation in snow densities measured using three different measurement methods from
November 27, 2015 to March 25, 2016.

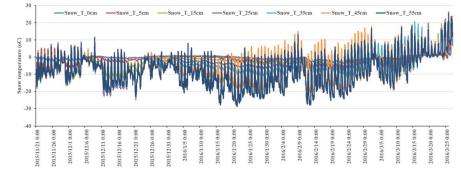
(b)



# Searth System Discussion Science Science Stressions

# 361362 **4.1.3 Snow temperature**

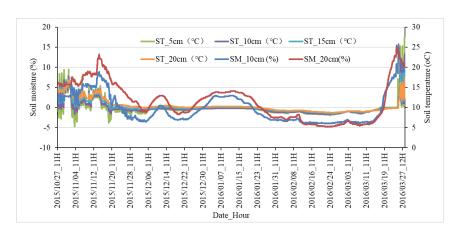
Snow temperature at 0 cm (snow/soil interface temperature) showed little diurnal variation, remaining at approximately -2.0 to 0.7°C. Snow temperature at top layer exhibited largest diurnal variation. The diurnal variation range decreased from top to bottom layers, and with the increase of snow depth, temperatures in more layers presented small diurnal variations (Figure 9). After March 17, 2016, the snow temperature of all layers were over 0°C which means snow cover did not refreeze anymore.



369 Figure 9: Variation in layering snow temperatures during experiment time.

# 370 4.2 Soil temperature and moisture

The soil temperature at 5 and 10 cm remained below 0 °C and stable during the snow season, but presented large fluctuation before snow cover onset and after snow off (Figure 10). The temperature gaps between 5 cm and 10 cm were much larger before snow cover onset than those during snow cover duration. The soil moistures at 10 cm were above 10% before snow cover onset and after snow off, and were below 10% during the snow cover duration. During Dec 12-14, and Jan 1- 20, soil moisture showed peak value, which corresponded to the two high-value periods of soil temperature.



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Figure 10: Hourly variation in soil temperature at 5 cm,10 cm, 15cm and 20 cm, and soil moisture at 10cm
and 20cm.



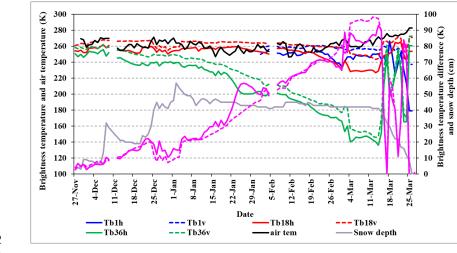


# 381 4.3 Brightness temperature

382 The microwave brightness temperatures varied with snow and soil characteristics, and weather 383 conditions. Figure 11a shows the daily brightness temperatures, brightness temperature difference 384 between 18 and 36 GHz, snow depth and air temperature at 1:00 am, and Figure 11b shows the hourly 385 variation in brightness temperatures at the three frequencies and air temperature after February 1. Data depict that Tb36h and Tb36v decreased during the whole snow season, Tb18h show obvious decline after 386 387 Feb 18, and Tb18v show decline after Mar 3 for vertical polarization (Figure 11a). After Jan 4, snow 388 depth stopped increasing, but the brightness temperature continued to decrease and brightness 389 temperature difference increased. Based on Figure 8, snow density arrived at stable on Jan 15. Therefore, 390 after Jan 4, the decreasing brightness temperatures was mainly caused by growing grain size.

After Feb 25, brightness temperature exhibited abrupt increase (at day time) - decrease (at night time) circle (Figure 11b), because air temperature at noon increasing up to above 270 K resulted in large liquid water content at day time, and the melted snowpack refroze when air temperature decreased at night time and brightness temperature decreased. After March 14, air temperature presented another big rise and even the night air temperatures were above 270 K; snow cover accelerated melting, and the liquid water could not be completely refrozen; then brightness temperature and brightness temperature difference exhibited unregularly variations.

The variation of L band was mainly influenced by soil moisture and soil temperature. We have soil temperatures at 0cm, 5 cm and 10 cm and soil moisture at 10 cm. However, the L band reflects the soil moisture within 5 cm which was absent in this experiment. Actually, we did not find the variation of brightness temperature at L band had relationship with soil moisture at 10 cm and soil temperature.







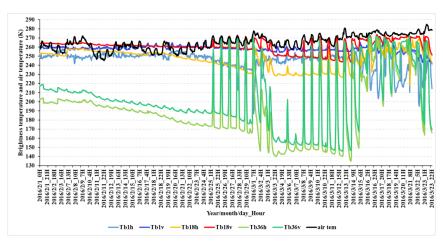
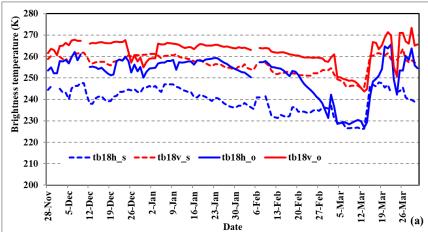


Figure 11: (a) Daily variations in brightness temperatures at 1.4 GHz, 18 GHz and 36 GHz, for horizontal
(Tb1h, Tb18h, Tb36h) and vertical polarizations (Tb1v, Tb18v, Tb36v), and the differences between Tb18h
and Tb36h (TBDh), and between Tb18v and Tb36v (TBDv), at 1:00 am (local time), from November 27,
2015 to March 26, 2016. (b) hourly variation in Tb1h, Tb18h, Tb36h, Tb1v, Tb18v, Tb36v, from February 1
to March 23, 2016.

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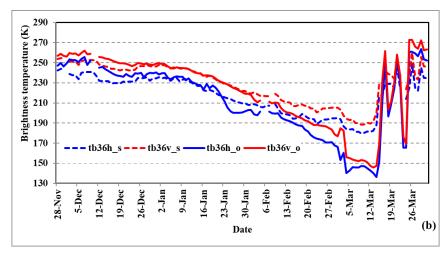
404

The brightness temperatures from AMSR-2 were compared with the ground-based observation at the overpass time (Figure 12). Although there was large difference between them, the general variations are the same, even for the abrupt change between Mar 3 and Mar 4, and the correlation coefficients at both polarizations were approximately 0.96 and 0.7 for 36 GHz and 18.7GHz, respectively. Satellite observed brightness temperature presented less decrease trend than ground-based observation, and the difference at 36 GHz is larger than at 18 GHz (Figure 12). The difference between ground-based and satellite observation might be attributed to the different viewing area.









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420 421

Figure 12: Comparison of brightness temperature between ground-based and Satellite-based observation (s: satellite; o: observation), (a) for 18 GHz, (b) for 36 GHz.

# 422 4.4 4-component Radiation

The land surface albedo is strongly related to the land cover. In this experiment, the down-short wave radiation presented general increase after January, and the trend became distinctive after February (Figure 13). The upward short-wave radiation abruptly increased when the ground was covered by snow, and sharply declined on the snow off day. From the first snowfall to the end of January, the ratios between upward and downward short-wave radiation were approximately 95%. The ratio decreased with snow age, and in the end of snow season the ratios decreased to below 50% because of increasing melted water.

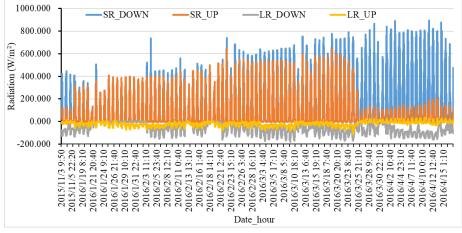


Figure 13: Minutely variation in 4-component radiation at Altay station from November 3 2015 to April 15



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# 432 5 Discussion

433 This experiment provided a suite of snow characteristics and microwave brightness temperatures, 434 and have proven useful for evaluating and updating microwave emission transfer model of snowpack 435 (Dai et al., 2021). This dataset reflected the general fact that brightness temperature at higher frequencies 436 presented stronger volume scattering of snow grains, and were more sensitive to snow characteristics. 437 This experiment revealed that the dominant control factor for the variation of brightness temperature was 438 the variation of grain size but not the snow depth. The largest snow depth or SWE does not correspond 439 to the largest brightness temperature gradient in the condition of dry snowpack. Due to the growth of 440 grain size, the peak gradient occurred before melting for stable snow cover. Therefore, the daily snow 441 depth variations curve derived from passive microwave remote sensing datasets tend to exhibit a 442 temporal offset from those of in situ observation.

443 During the snow season, brightness temperatures for both polarizations presented similar variation, 444 but they behaved different in some temporal periods. The horizontal polarization was more sensitive to 445 environment and demonstrated less stable than vertical polarization. Besides, the polarization difference at 18 GHz and 36 GHz showed increase and decrease trends, respectively during the experimental period. 446 447 The results for 18 GHz were opposite to the simulation results (Dai et al., 2021). These phenomena must 448 rely on the environmental conditions, snow characteristics and soil conditions. However, the subsurface 449 soil wetness data were absent, the dynamic ground emissivity could not be estimated. L band has strong 450 penetrability, and the brightness temperature variations were dominantly related to subsurface soil 451 condition, except when the liquid water content within snowpack was high. Therefore, although we did 452 not have soil moisture data in the subsurface layer, L band brightness temperatures were expected to retrieve soil moisture variation which influence the soil transmissivity (Babaeian et al., 2019; Naderpour 453 454 et al., 2017; Hirahara et al., 2020).

455 Snow surface albedo significantly influence the incoming solar radiation, playing an important role 456 in climate system. The factors to change the snow surface albedo essentially contains the snow characteristics (grain size, SWE, liquid water content, impurities, surface temperature etc), external 457 458 atmospheric condition and solar zenith angle (Aoki et al., 2003). Snow albedo was estimated based on 459 snow surface temperatures in some models (Roesch et al., 1999), and some models considered snow 460 surface albedo mainly depends on snow aging (Mabuchi et al., 1997). In this experiment, we obtained 461 the 4-component radiation, snow pit and meteorological data. These data provide nearly all observations 462 of possible influence factors, and could be utilized to discuss and analyze shortwave radiation process of 463 snowpack, and validate or improve multiple-snow-layer albedo model.

464 Snow grain sizes and snow densities within different layers presented different growth rates at 465 different temporal phase. Generally, the growth rates are related to the air temperature, pressure and snow 466 depth (Chen et al., 2020; Essery, 2015; Vionnet et al., 2012; Lehning et al., 2002); therefore, this dataset 467 can be used to analyze the evolution process of snow characteristics, as well as validation data for snow 468 models.

# 469 6 Conclusions

In a summary, the IMCS campaign provides a time series of snow pits observation, meteorological
 parameters, optical radiation and passive microwave brightness temperatures in the whole snow season
 of 2015/2016. The dataset is unique in providing microwave brightness temperatures and matched daily





473 snow pits data over a snow season at the fix site.

474 The dataset contains the unique daily snow pit data which present detail description of snow grain 475 size, snow density and snow temperature profiles. It can be used to analyzes the evolution process of 476 snow characteristics, validating or improving the snow process models, such as SNOWPACK (Lehning 477 et al., 2002), SNTHERM (Chen et al., 2020). The improvement of these models can further enhance the 478 prediction accuracy of land surface process and hydrology models, and the simulation accuracy of snow 479 microwave emission models. 480 Actually, this dataset has been utilized to analyze the volume scattering features of snow pack at 481 different frequencies (Dai et al., 2021). It can also be used to further analyze polarization characteristics of snow pack, and be used to validate different microwave emission models of snowpack. 482 483 Moreover, in this experiment, the microwave and optical radiations were simultaneously observed. 484 The existing studies reported that the optical equivalent diameter must be used in microwave emission 485 model with caution (Lowe and Picard, 2015; Roy et al., 2013). It is a good chance to analyze the 486 difference between the influence of grain size on microwave and optical radiation, establishing the bridge

487 between effective optical grain size and microwave grain size.

# 488 **7 Data availability**

489	The IMCS consolidated datasets are available after registration on the National Tibetan Plateau Data
490	Center and available online at https://data.tpdc.ac.cn/en/ (doi: 10.11888/Snow.tpdc.270886). Microwave
491	radiometry raw Data are available for scientific use on request from Northwest Institute of Eco-
492	Environment and Resources, Chinese Academy of Sciences.
493	
494	
495	Author contributions: LD and TC designed the experiment. LD, YZ, JT, MA, LX, SZ, YY YH and LX
496	collected the passive microwave and snow pit data. HL provided the 4-component radiation and snow
497	temperature data. LW provided meteorological data. LD write the manuscript, and TC made revision. All
498	authors contributed to the data consolidation.
499	
500	Competing interests: The authors declare that they have no conflict of interest.
501	
502	Acknowledgment: The authors would like to thank the Altay meteorological station for providing
503	logistics service and meteorological data.
504	
505	Financial support: This research was funded by the National Science Fund for Distinguished Young
506	Scholars (grant nos: 42125604), National Natural Science Foundation of China (grant nos: 42171143),
507	and CAS 'Light of West China' Program.
508	
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# 510 Appendix

0001表层	0002表层	0003表层	0004表层	0005表层	0006表层	0007表层	0008表层	0009表层	0010表层	0011表层	0012表层标尺
	1	entre l				h	24	U.S.*		2	
0013表层底标尺	0015表层底标尺	0016表层底标尺	0017表层底际尺	0018表层底标尺 0030 40 cm	0019表层底际尺 0031 40 cm	0020表层底标尺 0033 40 cm	0021_4th layer	0022_4th layer	0023_4th layer	0024_4th layer	0025_4th layer
			reference			644		Č.		<b>1</b>	
0040_40 cm reference	0041_30 cm	0042_30 cm	0043_30 cm	0044_30 cm	0046_30 cm	0048_30 cm	0049_30 cm reference	0050_20 cm	0052_20 cm	0053_20 cm	0054_20 cm
0055_20 cm	0056_20 cm	0057_20 cm	0058_20 cm reference	0059_15 cm	0060_15 cm	0061_15 cm	0062_15 cm	0063_15 cm	0064_15 cm	0065_15 cm reference	0066_8 cm
0067_8 cm	0068_8 cm	0069_8 cm	0070_8 cm	0071_8 cm	0072_8 cm	0073_8 cm	0074_8 cm reference	0075_1st	0076_1st	0077_1st	0078_1st

511

512 Figure A1: Photos of grains and reference ruler in each layer on February 15, 2016, and in each photo the

513 longest and shortest axis lengths of the chosen grains are labeled.

514

# 515 Table A1. Recorded longest and shortest axis length in Figure A.

Strati graph y	Thickn ess (cm)	Shape	Grain size (longest axis * shortest axis)(mm)								
the 3cm fifth	3cm	#22	0.595 *0.43 6	0.472 *0.47 1	0.450 *0.43 6	0.615 *0.47 4	0.374 *0.31 4	0.647 *0.30 7	0.656 *0.52 9	0.544 *0.51 9	0.717 *0.44 7
			0.750 *0.44 5	1.056 *0.95 5	0.623 *0.37 8	0.451 *0.40 5	1.397 *0.63 5	1.235 *0.32 7	0.600 *0.42 1	0.633 *0.55 6	0.729 *0.42 3
	3cm	3cm #37	2.605 *2.01 1	1.850 *1.32 8	1.626 *1.55 4	1.767 *1.68 5	1.718 *1.53 5	2.255 *1.29 6	1.674 *1.60 1	1.542 *1.26 9	3.505 *1.44 0
the fourth			3.055 *1.77 4	1.448 *1.37	2.461 *1.91 4	2.757 *2.11 5	2.179 *2.05 9	2.393 *1.78 8			
the third	25cm	#27, #31, #37	2.569 *1.60 7	2.073 *2.13 0	2.591 *1.41 4	1.869 *1.80 2	2.067 *1.26 6	1.209 *1.10 6	1.719 *1.18 8	1.648 *0.97 5	1.911 *1.58 2





	ľ										
			1.921	1.518	1.291	1.690	1.756	1.812	1.733	1.880	2.411
			*1.71	*1.06	*1.14	*1.55	*1.39	*1.26	*1.67	*1.51	*1.22
			0	7	7	1	8	3	2	8	0
			2.118	1.614	1.795	2.215	1.864	1.967	2.008	1.362	1.484
			*1.72	*1.45	*1.70	*2.31	*1.69	*1.65	*1.39	*1.14	*1.29
			7	7	5	1	2	1	5	1	1
			4.251	3.012	2.805	1.799	1.402	3.040	2.850		
		#33, #34	*2.26		*1.99	*1.41	*1.19	*2.07	*2.09		
the	12		6	*2.65	5	5	5	3	5		
secon d	12		3.900	2.420	2.515	2.044	2.506	2.894	2.413	2.494	4.929
u			*2.53	*2.33	*2.20	*2.03	*2.36	*2.16	*1.95	*1.81	*3.25
			2	3	6	2	3	1	0	6	7
		#40	4.933	3.207	3.562	2.818	3.581	6.179			
the		#40,	*3.37	*2.77	*1.70	*1.66	*2.51	*3.56			
first	4	#34,	8	4	1	8	8	2			
		#38									

516

517

# 518 Table A2: One example of record table for snow density observation.

observation date:	20160111	observation	19:03-9:40	weather:	clear	snow depth:	48cm
5	now Folk table				snow tube	table	
observation height (cm)	liquid water content(%)	snow density (g/cm3)		snow depth(cm)	46.5	47	47.5
	0	0.1923	1	snow pressure(g/cm2)	9.1	9	9.5
5	0.118	0.1882		snow density(g/cm3)	0. 1957	0.1915	0.2000
	0	0.1882					
	0.461	0.164			snow shovel	table	
10	0.46	0.1631		observation layer (cm)	weight of shovel+snow(g)	weight of shovel(g)	snow density(g/cm3
	0.461	0.1361			865.04	572.16	0. 1953
	0.123	0.2532		0-10	858.72	572.16	0.1910
15	0	0.2506			866. 69	572.16	0.1964
	0	0.2417			878.58	572.16	0.2043
	0.24	0.2159		10-20	887.04	572.16	0.2099
20	0.119	0.2155			872.79	572.16	0.2004
	0.119	0.2146			905.34	572.16	0.2221
25	0.117	0.1977	1	20-30	903. 41	572.16	0. 2208
	0	0.1994			907.88	572.16	0. 2238
	0	0.1984			832.75	572.16	0.1737
	0	0.1919		30-40	838.14	572.16	0.1773
30	0	0.1966			837.27	572.16	0.1767
	0	0.1928					
	0	0.1534	1	40-50			
35	0	0.1517					
	0	0.1472					
	0.325	0.1097	1	50-60			
40	0	0.1054					
	0.107	0.1088					
	0	0.0922					
45	0	0.0991					
	0	0.0928					
50							
55		_					

519 520





# 522 References:

- 523 Babaeian, E., Sadeghi, M., Jones, S.B., Montzka, C., Vereecken, H., and Tuller, M.: Ground, Proximal,
- 524 and Satellite Remote Sensing of Soil Moisture. Reviews of Geophysics, 57(2), 530-616, doi:
- 525 10.1029/2018RG000618, 2019.
- 526 Barnett, T.P., Adam, J.C., and Lettenmaier, D.P.: Potential impacts of a warming climate on water
- 527 availability in snow-dominated regions. *Nature*, 438, 303-309, doi: 10.1038/nature04141, 2005.
- 528 Brucker, L., Hiemstra, C., Marshall, H.-P., Elder, K., De Roo, R., Mousavi, M., Bliven, F., Peterson,
- 529 W., Deems, J., Gadomski, P., Gelvin, A., Spaete, L., Barnhart, T., Brandt, T., Burkhart, J., Crawford,
- 530 C., Datta, T., Erikstrod, H., Glenn, N., Hale, K., Holben, B., Houser, P., Jennings, K., Kelly, R., Kraft,
- 531 J., Langlois, A., McGrath, D., Merriman, C., Molotch, N., Nolin, A., Polashenski, C., Raleigh, M.,
- 532 Rittger, K., Rodriguez, C., Roy, A., Skiles, M., Small, E., Tedesco, M., Tennant, C., Thompson, A.,
- 533 Tian, L., Uhlmann, Z., Webb, R., Wingo, M., and Ieee: A FIRST OVERVIEW OF SNOWEX
- 534 GROUND-BASED REMOTE SENSING ACTIVITIES DURING THE WINTER 2016-2017. 2017
- 535 Ieee International Geoscience and Remote Sensing Symposium (pp. 1391-1394), 2017
- 536 Che, T., Dai, L.Y., Zheng, X.M., Li, X.F., and Zhao, K.: Estimation of snow depth from passive
- 537 microwave brightness temperature data in forest regions of northeast China. Remote Sensing of
- 538 Environment, 183, 334-349, doi: 10.1016/j.rse.2016.06.005, 2016.
- 539 Che, T., Li, X., Jin, R., Armstrong, and R., Zhang ,T.J. : Snow depth derived from passive microwave
- remote-sensing data in China. *Annals of Glaciology*, 49, 145. doi: 10.3189/172756408787814690,
  2008.
- 542 Chen, T., Pan, J.M., Chang, S.L., Xiong, C., Shi, J.C., Liu, M.Y., Che, T., Wang, L.F., and Liu, H.R.:
- 543 Validation of the SNTHERM Model Applied for Snow Depth, Grain Size, and Brightness Temperature
- 544 Simulation at Meteorological Stations in China. Remote Sensing, 12, 507, doi: Artn
- 545 50710.3390/Rs12030507, 2020.
- 546 Cline, D., Elder, K., Davis, B., Hardy, J., Liston, G., Imel, D., Yueh, S., Gasiewski, A., Koh, G.,
- 547 Amstrong, R., and Parsons, M.: An overview of the NASA Cold Land Processes Field Experiment
- 548 (CLPX-2002). Microwave Remote Sensing of the Atmosphere and Environment Iii, 4894, 361-372.
- 549 doi: Doi 10.1117/12.467766, 2003.
- 550 Cohen, J: Snow cover and climate. Weather, 49, 150-156, 1994.
- 551 Dai, L. (2020): Microwave radiometry experiment data in Altay (2015/2016). National Tibetan Plateau
- 552 Data Center [dataset]. doi: 10.11888/Snow.tpdc.270886, 2020.
- 553 Dai, L.Y., Che, T., Wang, J., and Zhang, P. :Snow depth and snow water equivalent estimation from
- 554 AMSR-E data based on a priori snow characteristics in Xinjiang, China. Remote Sensing of
- 555 Environment, 127, 14-29,. doi: 10.1016/j.rse.2011.08.029, 2012.
- 556 Derksen, C., Toose, P., Lemmetyinen, J., Pulliainen, J., Langlois, A., Rutter, N., and Fuller, M.C.:
- 557 Evaluation of passive microwave brightness temperature simulations and snow water equivalent
- 558 retrievals through a winter season. Remote Sensing of Environment, 117, 236-248, doi:
- 559 10.1016/j.rse.2011.09.021, 2012.
- 560 Ding, Y.J., Yang, J.P., Wang, S.X., and Chang, Y.P.: A review of the interaction between the
- 561 cryosphere and atmosphere. Sciences in Cold and Arid Regions, 12 (6): 329-342, doi:
- 562 10.3724/SP.J.1226.2020.00329, 2020.
- 563 Essery, R.: A factorial snowpack model (FSM 1.0). Geosci. Model Dev. 2015, 8, 3867–3876.





- 564 Hirahara, Y., de Rosnay, P., and Arduini, G.: Evaluation of a Microwave Emissivity Module for Snow
- 565 Covered Area with CMEM in the ECMWF Integrated Forecasting System. Remote Sensing, 12(18),
- 566 doi: Artn 294610.3390/Rs12182946, 2020.
- 567 Immerzeel, W.W., van Beek, L.P.H., and Bierkens, M.F.P.: Climate Change Will Affect the Asian
- 568 Water Towers. Science, 328(5984), 1382-1385. doi: 10.1126/science.1183188, 2010.
- 569 Jiang, L.M., Wang, P., Zhang, L.X., Yang, H., and Yang, J.T.: Improvement of snow depth retrieval for
- 570 FY3B-MWRI in China. Science China-Earth Sciences, 57, 1278-1292, doi: 10.1007/s11430-013-4798-
- 571 8, 2014.
- 572 Jordan, R.E.: A One-Dimensional Temperature Model for a Snow Cover: Technical Documentation for
- 573 SNTHERM.89; U.S. Army Cold Regions Research and Engineering Laboratory: Hanover, NH, USA,574 1991.
- 575 Lehning, M., Bartelt, P., Brown, B., Fierz, C., and Satyawali, P.: A physical SNOWPACK model for
- the Swiss avalanche warning Part II: Snow microstructure. *Cold Regions Science and Technology*, *35*,
  147-167, Doi 10.1016/S0165-232x(02)00073-3, 2002.
- 578 Lemmetyinen, J., Kontu, A., Pulliainen, J., Vehvilainen, J., Rautiainen, K., Wiesmann, A., Matzler, C.,
- 579 Werner, C., Rott, H., Nagler, T., Schneebeli, M., Proksch, M., Schuttemeyer, D., Kern, M., and
- 580 Davidson, M.W.J. : Nordic Snow Radar Experiment. Geoscientific Instrumentation Methods and Data
- 581 Systems, 5, 403-415, doi: 10.5194/gi-5-403-2016, 2016.
- 582 Löwe H. and Picard, G. "Microwave scattering coefficient of snow in MEMLS and DMRT-ML
- 583 revisited: The relevance of sticky hard spheres and tomography-based estimates of stickiness,"
- 584 Cryosphere, vol. 9, no. 6, pp. 2101–2117, Nov. 2015.
- 585 Mortimer, C., Mudryk, L., Derksen, C., Luojus, K., Brown, R., Kelly, R., and Tedesco, M. : Evaluation
- 586 of long-term Northern Hemisphere snow water equivalent products. Cryosphere, 14(5), 1579-1594,
- 587 doi: 10.5194/tc-14-1579-2020, 2020.
- 588 Naderpour, R., Schwank, M., Matzler, C., Lemmetyinen, J., and Steffen, K.: Snow Density and Ground
- 589 Permittivity Retrieved From L-Band Radiometry: A Retrieval Sensitivity Analysis. Ieee Journal of
- 590 Selected Topics in Applied Earth Observations and Remote Sensing, 10(7), 3148-3161, doi:
- 591 10.1109/Jstars.2017.2669336, 2017.
- 592 Pulliainen, J., Luojus, K., Derksen, C., Mudryk, L., Lemmetyinen, J., Salminen, M., Ikonen, J., Takala,
- 593 M., Cohen, J., Smolander, T., and Norberg, J.: Patterns and trends of Northern Hemisphere snow mass
- 594 from 1980 to 2018. Nature, 581(7808), 294-298. doi: 10.1038/s41586-020-2258-0, 2020.
- 595 Roy, A., Picard, G., Royer, A., Montpetit, B., Dupont, F., Langlois, A., Derksen, C., and Champollion,
- 596 N.: Brightness Temperature Simulations of the Canadian Seasonal Snowpack Driven by Measurements
- 597 of the Snow Specific Surface Area. Ieee Transactions on Geoscience and Remote Sensing, 51, 4692-
- 598 4704, doi: 10.1109/Tgrs.2012.2235842, 2013.
- 599 Roy, A., Picard, G., Royer, A., Montpetit, B., Dupont, F., Langlois, A., Derksen, C., and Champollion,
- 600 N.: Brightness Temperature Simulations of the Canadian Seasonal Snowpack Driven by Measurements
- 601 of the Snow Specific Surface Area, IEEE T. Geosci. Remote, 51, 4692–4704,
- 602 doi:10.1109/TGRS.2012.2235842, 2013
- 603 Tedesco, M., Narvekar, P.S.: Assessment of the NASA AMSR-E SWE Product. Ieee Journal of Selected
- Topics in Applied Earth Observations and Remote Sensing, 3, 141-159, doi:
- 605 10.1109/Jstars.2010.2040462, 2010.
- 606 Vionnet, V., Brun, E., Morin, S., Boone, A., Faroux, S., Le Moigne, P., Martin, E., and Willemet, J.M.:
- 607 The detailed snowpack scheme Crocus and its implementation in SURFEX v7.2. Geoscientific Model





- 608 Development, 5, 773-791, doi: 10.5194/gmd-5-773-2012, 2012.
- 609 Xiao, L., Che, T., and Dai, L.Y.: Evaluation of Remote Sensing and Reanalysis Snow Depth Datasets
- 610 over the Northern Hemisphere during 1980-2016. Remote Sensing, 12(19), doi: Artn
- 611 325310.3390/Rs12193253, 2020.
- 612 Yang, Z.L., Dickinson, R.E., Robock, A., and Vinnikov, K.Y.: Validation of the snow submodel of the
- 613 biosphere-atmosphere transfer scheme with Russian snow cover and meteorological observational data.
- 614 Journal of Climate, 10, 353-373, doi: 10.1175/1520-0442(1997)010<0353:Votsso>2.0.Co;2, 1997.