1 Microwave radiometry experiment for snow in Altay

China: time series of *in situ* data for electromagnetic and physical features of snow pack

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Abstract, Snow depth is a key parameter in climatic and hydrological systems, Passive microwave 19 20 remote sensing, snow process model and data assimilation are the main methods to estimate snow depth 21 in large scale. The estimation accuracies strongly depend on input of snow parameters or characteristics. 22 snow depth Because the evolving processes of snow parameters vary spatiotemporally, and are difficult 23 to accurately simulate or observe, large uncertainties and inconsistence exist among existing snow depth products. Therefore, a comprehensive experiment is needed to understand the evolution processes of 24 25 snow characteristics and their influence on microwave radiation of snowpack, to evaluate and improve 26 the snow depth and SWE retrieval and simulation methods. In this paper, we present a comprehensive 27 experiment, namely, An-Integrated Microwave Radiometry Campaign for snow (IMCS), was conducted 28 at the Altay National Reference Meteorological station (ANRMS) _ in Xinjiang, China, during snow 29 season of 2015/2016. The campaign hosted a dual polarized microwave radiometer operating at L, K and 30 Ka bands to provide minutely passive microwave observations of snow cover at a fixed site, along with 31 daily manual snow pit observation of snow physical parametersmeasurements, automatic observation of 32 ten-minute 4-component radiation and layered snow temperatures, and meteorological observation of hourly weather data and soil dataten minute automatic 4 component radiation and layered snow 33 temperatures, covering a full snow season of 2015/2016. The measurements of meteorological and 34 35 underlying soil parameters were requested from the ANRMS. This study provides a summary of the 36 obtained data, detailing measurement protocols for microwave radiometry, in situ snow pit and station 37 observation data. A brief analysis of the microwave signatures against snow parameters is presented. 38 These data were consolidated into five NetCDF files and A consolidated dataset of observations, 39 comprising the ground passive microwave brightness temperatures, in situ snow characteristics, 4component radiation and weather parameters, was achieved at the National Tibetan Plateau Data Center, 40

41 China. To the best of our knowledge, our The dataset is unique in providing continuous daily snow pits 42 data and coincident microwave brightness temperatures, radiation and meteorological data, at a fixed site 43 over a full season..., which can be straightforwardly used for The dataset is expected to serve the 44 evaluation and development of microwave radiative transfer models and snow process models., along 45 for land surface process and hydrology models. The consolidated data are available at 46 http://data.tpdc.ac.cn/zh-hans/data/df1b5edb-daf7-421f-b326-cdb278547eb5/ (doi: 47 10.11888/Snow.tpdc.270886) (Dai, 2020).

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49 Key words: Snow, Microwave radiometry, Snow pit, Experiment

50 1 Introduction

51 Seasonal snow cover plays a critical role in climate and hydrological systems (Cohen, 1994; Ding 52 et al., 2020; Barnett et al., 2005; Immerzeel et al., 2010) by its high albedo, thermal insulation, fresh 53 water reserves and its phase change processes. Snow cover can be accurately identified by optical remote 54 sensing. However, the snow surface albedo is controlled by snow characteristics (Aoki et al., 2003 and 55 2000), and variations in snow characteristics cause uncertainties in albedo estimation. Snow depth and snow water equivalent (SWE) are currently estimated using passive microwave at global and regional 56 57 scales (Pullianen et al., 2020; Tedesco and Narvekar, 2010; Jiang et al., 2014; Che et al., 2008). Although 58 several global and regional snow depth and SWE products have been released, large uncertainties exist 59 in these products because of the spatio temporal variations in snow characteristics (Xiao et al., 2020; 60 Mortimer et al., 2020; Che et al., 2016; Dai et al., 2012; Dai and Che, 2022). Therefore, the observation of electromagnetic and physical parameters of snowpack are necessary to improve understanding of the 61 electromagnetic radiation process of snowpack to enhance the estimation accuracy of snow surface 62 63 albedo and snow depth.

64 Field experiments/campaigns, as the main and most important approach for snow studies, To 65 evaluate and improve snow depth and SWE retrieval methods from passive microwave remote sensing 66 observations and to combine remote sensing technologies with modeling and data assimilation methods 67 to produce the most accurate products, a few large or systematic experiments or campaigns-have been 68 conducted to obtain the <u>on</u> electromagnetic and physical characteristics measurement of snow cover. 69 The mainse experiments/campaigns are summarized in table 1. The Cold Land Processes Field 70 Experiment (CLPX) (https://nside.org/data/clpx/index.html), one of the most well-known experiments, 71 was carried out from winter of 2002 to spring of 2003 in Colorado, USA (Cline et al., 2003). During the 72 campaign, snow pits were collected in February and March of 2002 and 2003 to coincide with airborne 73 and ground remote sensing observations. <u>In 2017, NASA</u> SnowEx campaign 74 (https://nsidc.org/data/snowex) was conducted in 2017-in ColaradoColorado to test and develop 75 algorithms for measurement of SWE in forested and non-forested areas by providing multi-sensor 76 observations of seasonally snow-covered landscapes (Brucker et al., 2017). The campaign is still ongoing 77 and will be conducted in other areas with different snow conditions. In northern Canada, mobile sled-78 mounted microwave radiometers were deployed in forest, open and lake environments from November 79 2009 to April 2010 and snow characteristics within the footprints of radiometers were measured to 80 improve the understanding of the influence of snow characteristics on brightness temperatures (Derksen 81 et al., 2012; Roy et al., 2013). These aforementioned microwave experiments were of mobile observation. 82 The observations were conducted at different In these experiments, there were multiple observation sites

for different land cover, resulting in good representativity for evaluating snow microwave emission model (Tedesco and Kim, 2006; Royer et al., 2017), however, with relative short temporal range, while dense temporal resolution is important to reveal the evolution of snow characteristicsbut relative short temporal range. The snow pit observations were used toaimed at evaluate evaluating snow microwave emission model in different land cover (Tedesco and Kim, 2006; Royer et al., 2017), but they did could not exhibit the evolution of snow parameters.

89 In the Arctic region, the Nordic Snow Radar Experiment (NoSREx) campaign was conducted at a 90 fixed field of the Arctic boreal forest area in Sodankylä, Finland, during 2009 ~ 2013 (Lemmetyinen et 91 al., 2016). Theis experiment provided a -continuous time series of active and passive microwave 92 observations of snow cover at a representative location of the Arctic boreal forest area spanning an entire 93 winter season, with synchronous observations of-and matched snow pit observations were made weekly. 94 In Asia, an experiment of radiation budget over snow cover (JERBES) was conducted in Japan. In the 95 experiment, snow pit work at 3 or 4 day intervals was conducted simultaneously with radiation budget observations during winter of 1999/2000 and 2000/2001 to analyze the effects of snow physical 96 97 parameters on albedo (Aoki et al., 2003). The NoSREx and JERBES Japan radiation experiments, were 98 for fixed field observation, which provided longer-improved time series of data than CLPX and SnowEx. 99 These experiments were conducted in deep snow areas, and the wWeekly observation could reflect 100 general evolution process of snow characteristics but might miss some key details that occur at sub-101 weekly scales. In the Tibetan plateau with shallow snow cover, multiple years of microwave radiometry 102 observation at L band were conducted to study passive microwave remote sensing of frozen soil (Zheng 103 et al., 2019, <u>and</u> 2021a and; Zhang et al., 2021b). However, in the long-term series of experiment, no 104 snow pit was measured and the microwave radiometry observation was performed at L band which is 105 insensitive to snowpack.

Table 1. Summary of existing experiments for microwave and optical radiation and physical features of snowpack

Campaign	Location	Temporal range	Observation content
CLPX	Different sites in Colorado,	February and March of 2002 and 2003	Inconsecutive multiple sensor observation, including microwave radiometry over snow, <u>and</u> and matchedsynchronous snow pit measurements were conducted at different sites with short temporal range
SnowEx-year 1	Grand Mesa, and Senator Beck Basin, Colorado	February of 2017	Inconsecutive multiple sensor observation, including microwave radiometry over snow, and matched_synchronous_snow pit measurements were conducted_at different sites with short temporal range-
CMRES ¹	Mobile observation at Forest, open and lake in the northern Canadian region	November of 2009-April of 2010	Mobile microwave radiometry and snow pit observation within footprint of radiometer. Short temporal range and inconsecutive observation

NoSREx	Fixed site in	Snow season	Consecutive microwave radiometry and SAR
	Sodankylä, Finland	during 2009-	observation over snow, and weekly snow pit
		2013	measurement
JERBS ²	Fixed site in Japan	Snow season	Consecutive optical radiation observation over
		during 1999-	snow and consecutive snow pit measurement at 3
		2000	or 4-day interval .
IMCS_	Fixed site in China	November of	Consecutive microwave radiometry and optical
(Presented in		2015-March of	radiation observation, and consecutive daily snow
<u>this work)</u>		2016	pit measurements .

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Note: ¹CMRES: Microwave radiometry experiment on snow cover conducted in northern Canada ²JERBS: Experiment of radiation budget over snow cover in Japan

111 To understand the evolution of snow characteristics and their influence on passive microwave 112 brightness temperatures and radiation budget, an integrated field experiment on snowpack was conducted during a full snow season, in Altay, China. The experiment was designed to cover periods from snow-113 114 free conditions to eventual snow melt-off during 2015/2016. The microwave radiometry measurements 115 at L, K and Ka bands for multiple angles were complemented by a dual-polarized microwave radiometer 116 with 4-component radiation and daily in situ observations of snow, soil and atmospheric properties, using 117 both manual and automated methods. To the best of our knowledge, our dataset is unique in providing 118 continuous daily snow pit observations over a full snow season at a fixed site, along with synchronous 119 microwave brightness temperatures, radiation and meteorological data. The dataset -of electromagnetic 120 and physical parameters wereis __further consolidated and organized, which can be to be easily read and 121 utilized used for other researchers with interests.

The dataset is unique in providing continuous daily snow pits data over a snow season at a fixed site and matched microwave brightness temperatures, radiation and meteorological data. In the next section, the experiment location, parameters, and parameter measurement protocols are described;-. section_Section_3 introduces the consolidated data which was released at the National Tibetan Plateau Data Center, China;-_section_4 presents content of brightness temperature, 4-component radiation, snow pit data, soil temperature and moisture, and meteorological data;-_section_Section_5 discusses the possible applications and uncertainties;-_and fFinally the conclusions are summarized in section 6.

129 2 Description of experiment setup

130 2.1 Measurement location

The Integrated Microwave Radiometry Campaign for snow (IMCS) was performed during the 2015/2016 snow season (from November 27, 2015 to March 25, 2016) at the Altay National Reference Meteorological station (ANRMS) (N47°44'26.58", E 88°4'21.55"), which is approximately 6 km from the foot of Altay mountain in the northwest China (Figure 1). Altay mountain with elevation up to 3000 m, running northwest and southeast, is at the junction of China, Russia, Mongolia and Kazakhstan, and providesing snow water resources for these four countries. The average annual maximum snow depth measured in this station is approximately 40 cm, with a maximum over 70 cm. In the southwest of Altay mountain, crop land and desert with flat terrain are the dominant land covers. Snow cover is critical fresh
water for the irrigation in this area. In this experiment, measurements included microwave radiometry,
4-component radiation, snow pit and soil parameters. The test sites of this experiment was-were four
neighboring bare rectangle fields in the ANRMS with areas of 2500_m² (black rectangle field in Figure 1), 2500_m² (pink rectangle field in Figure 1), 200_m² (red rectangle field in Figure 1) and 400 m² (blue
rectangle field in Figure 1), respectively.

In the pink field, the ground-based microwave radiometer was set up in the middle of the field, facing south to collect brightness temperatures <u>of over</u> snow cover. The black field behind the microwave radiometers (north of the radiometers) was for manual snow pit data collection. The microwave radiometer observations and snow pit data collections were conducted by Northwest Institute of Eco-Environment and Resources, Chinese Academy of Science (NIEER) from November 27, 2015 to March 25, 2016 (snow free aAfter March 25, 2016, snow melted out).

150 The blue field was for meteorological measurements including wind speed, wind direction, air 151 temperature, air wetness, air pressure, precipitation, soil temperature, soil moisture among others. These 152 parameters were automatically obtained from instruments, and the instruments setup and data collection 153 were operated by ANRMS. This station also has daily manual observation of snow depth and SWE. In 154 this experiment, we requested the wind, air pressure, air wetness, air pressure, soil temperature and 155 moisture data during this experiment from ANRMS. The red field was designed for automatic 156 measurement of layered snow temperatures, snow density, SWE, snow depth, and albedo-, with These automatic measurement instruments were installed and maintained operated by NIEER, and started 157 158 working from since 2013. However, during the experiment, the instruments for snow density and SWE 159 did not workwere not functional, and we only collected layered snow temperatures and 4-component 160 radiation.

161 <u>Because Because the</u> four observation fields are located within the domain of the station, are and 162 the distance between them are with distance less than 100 m, it is reasonable to assume that the snow 163 characteristics and soil and weather conditions are thought to be the same consistent among these four 164 <u>fields</u>. Overall, the experiment performed a systematic observation covering electromagnetic and 165 physical features of snow pack, providing data for studies on snow remote sensing and models.



167 Figure 1+. Location of the Altay National Reference Meteorological station (ANRMS) in Asia, along with and 168 the distribution of the three-four test sites experiment fields in the ANRMS. The black rectangle field 169 (approximately 40 m × 50 m) represents the field used for snow characteristics (approximately 40 m × 50 m) 170 includingwas for snow layering, layer thickness, snow density, snow grain size and shape of each layer. The 171 <u>pink rectangle (approximately 60 m \times 50 m) was for, and microwave radiometers (approximately 60 m \times 50 m) was for, and microwave radiometers (approximately 60 m \times 50 m) was for, and microwave radiometers (approximately 60 m \times 50 m) was for, and microwave radiometers (approximately 60 m \times 50 m) was for, and microwave radiometers (approximately 60 m \times 50 m) was for, and microwave radiometers (approximately 60 m \times 50 m) was for, and microwave radiometers (approximately 60 m \times 50 m) was for, and microwave radiometers (approximately 60 m \times 50 m) was for, and microwave radiometers (approximately 60 m \times 50 m) was for, and microwave radiometers (approximately 60 m \times 50 m) was for, and microwave radiometers (approximately 60 m \times 50 m) was for the formeters (approximately 60 m \times 50 m \times 50 m) was for the formeters (appr</u> 172 m)-observations. The blue rectangle field is wasthe field __ for meteorological and soil data collection operated 173 by the ANRMS. The red rectangle field wasis for the automatically observation of the snow temperature, and 174 4-component radiation, designed by Northwest Institute of Eco-Environment and Resources, Chinese 175 Academy of Science (NIEER).

176 2.2 Measurement methods

177 The microwave signatures from snowpack vary with snow characteristics, soil and weather 178 conditions. In this experiment, the measurements include microwave radiometry observation to collect 179 brightness temperature, manual snow pit observation to collect snow physical parameters, automatic 180 observation to collect 4-component radiation and snow temperatures, and meteorological observation 181 which containscontaining weather data and soil data.

182 **2.2.1. Microwave radiometry**

183The brightness temperatures at 1.4 GHz, 18.6 GHz, and 36.5 GHz for both polarization (Tb1h, Tb1v,184Tb18h, Tb18v, Tb36h, Tb36v) were automatically collected using a six-channel dual polarized185microwave radiometer RPG-6CH-DP (Radiometer Physics GmbH, Germany,186https://www.radiometerphysics.de/products/microwave-remote-sensing-instruments/radiometers/). The

technical specifications of the RPG-6CH-DP are described in Table 2. The RPG-6CH-DP contains a
 built-in temperature sensor which canused for measureing air temperature. The automated data collection

- 189 frequency was set to 1 minute.
- 190

191 Table 2. Technical Specifications of the RPG-6CH -DP Microwave Radiometer.

Parameter	Value		
Manufacturer	Radiometer Physics GmbH		
System noise temperatures	<900 K		
Bandwidth	400MHz (20MHz for 1.4 GHz)		
System stability	0.5 K		
Dynamic range	0~400 K		
Frequencies (GHz)	1.4, 18.7, 36.5		
Polarizations	V, H		
Tetermel cellbration	Internal Dicke switch and software control for		
internal canoration	automatic sky tilt calibration		
Receiver and antenna thermal stabilization	< 0.015 K		
Antenna sidelobe level	< -30 dBc		
Optical resolution (HPBW)	6.1° (11° for 1.4 GHz)		
Incidence angle	0~90°		
Azimuth angle	360°		

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193 Before the snow season, a platform with height of 5 m, length of 4 m and width of 2 m was 194 constructed in the experiment field (Figure 2). A 4-m orbit was fixed on the platform. The RPG-6CH-DP 195 was set up on the orbit and could be moved along the orbit. The microwave radiometers at K and Ka 196 bands began working from November 27, 2015, but-while the L band radiometer did notbegan working 197 since until-January 30, 2016. These radiometers were sky tipping calibrated, and the calibration with 198 accuracy is of 1 K. In clear sky conditions, the sky brightness temperatures were approximately 29.7 ± 0.3 199 K at 18.7 GHz for both polarizations and 29.3 ± 0.9 K at 36.5 GHz for both polarizations. But While the 200 sky brightness temperatures at L band showed large fluctuation. They ranged from -1 to 8 K for horizontal polarization, and 1 to 16 K for vertical polarization. 201

202 Generally, the radiometers were fixed in the middle of the orbit to observe snow cover with incidence 203 angle of 50°. Multi-angle observations were conducted after every big snowfall, or and every 5 days in 204 the stable period. In the melt period, observation frequency increased. There are total seventeen multi-205 angle observation (on December 3, 19, and 30; January 3, 8, 13, 18, 3, and 28; February 3; March 3, 10, 206 15, 22, 26, 28, and 31,) when the radiometer was set to scan the ground at different incidence angles at 207 two ends and the middlethe left, middle and right of the orbit and the middle place of the orbit, 208 respectively. Although the view fields of the antennas for 1.4 GHz, 18.7 GHz and 36.5 GHz did not 209 completely overlap, the measured results showed that the brightness temperatures observed by 210 radiometers at the left, middle and right of the orbit varied less thanwithin 1 K. Therefore, the snow and 211 soil characteristics conditions were considered homogeneous within the view fields of the three antennas.



Figure 2. Ground-based microwave radiometer observation system.

2.2.2 Snow pit measurement

The snow characteristics were obtained by manual snow pit measurements in the black field, including snow layering, snow layer thickness, grain size, snow density, and snow temperature, were collected by manual snow pit measurements in the black fields. These data were daily collected during 8:00-10:00 am local time, from November 27, 2015 to March 25, 2016, except 7 days (please see Table 3). Although the snow temperatures were manually measured at snow pits, the automatically collected snow temperatures in the red field were utilized in this study, because the temperature measured at snow pits could not reflect the natural temperature profile when the snow pits exposed to air.

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Table 3. Variables collected by manual daily snow pit measurement in black field in figure 1, and along with their observation instruments, observation time and frequencies.

Parameter	Instruments	Precision	Layering style	Observati on time or frequency	Absent date
Layer thickness	Ruler	0.1cm	Natural		no
(cm)			layering	local time	
Snow density (g/cm ³)	Snow tube (Chinese Meteorological Administration)	pressure:0. 1g/cm ² , snow depth: 0.1 cm	Whole snowpack	8:00- 10:00 am	no
Snow density		weight: 0.01g,	Every 10		January 2-
(g/cm ³)	Snow shovel (NIEER)	volume: 1cm ³	cm		3, 2016;

Snow density (g/cm ³) and	Snow fork (Toikka Enginnering Ltd.)	0.0001g/cm ³	Every 5 cm	February 20, 2016
Liquid water content (%)	Snow fork	0.001%	Every 5 cm	
Snow grain size (mm)	Anyty V500IR/UV	0.001mm	Natural layering	December 24, 31,
Snow grain shape	Shape card	N/A	Natural layering	2015; January 1- 3, 23, 2016, February 20, 2016

227 The first step of snow pit measurement is making a snow pit. In the black field, a new snow pit was 228 dug each day using . As spade was used to excavate snow pit. The length of the snow pit profile was 229 approximately 2 m x 1m to make sure all parameters <u>could bewere</u> measured from unbroken snowpack. 230 The width of the snow pit was approximately 1 m. The snow pit section was made as flat as possible 231 using a flat shovel or ruler. When the snow profile is exposed to air for a long time, the snow 232 characteristics will be influenced by environment and will be different from the natural snow 233 characteristics. In order to make sure every observation conducted on natural snow pit, the snow pit was 234 backfilled with the shoveled snow after finishing all observations, and the new snow pit in the following 235 day was made at least 1-m distance away from the last-latest snow pit. After finishing a snow pit, the 236 natural snowpack stratification was then visually determined, and the thickness of each layer was 237 measured usingrecorded against a ruler.

238 The third-next step was measuring grain size and shape type in each layer. The grain size and type 239 within each natural layer were estimated visually from using a microscope with an "Anyty V500IR/UV" 240 camera (Figure 3a). A-The software "VIEWTER Plus" matched the microscope was used to measure 241 grain size. The grain type was determined based on Fierz et al. (2009). In this experiment, we utilized 242 the length of longest axes and the length of shortest axes to describe grain size (Figure 3b). When using 243 the software to measure the grain size, a reference must be needed. In this experiment, a ruler with marked 244 0.5 mm marking-was used as a reference (Figure 3c). We adjusted the focus of the camera to make sure 245 the grains at the clearest status in camera to take photos, and the photo of ruler scale was taken at the 246 same focus. If the thickness of one layer was less than 10 cm, measurements were performed at the top 247 and bottom of the layer, respectively. If the thickness was greater than 10 cm, measurements were 248 performed at the top, middle, and bottom of the layer, respectively. For each layer, at least 5 photos were 249 taken, and the longest axes length and the shortest axes length of at least 10 typical grains were chosen 250 to measure recorded the longest axes length and the shortest axes length in the photos of each layer. Each 251 layer had at least 10 groups of longest and shortest axes length; tThe final grain size was the average of 252 these values 10-group recorded values. Figure A1 presents an example of the original photos of grains in 253 each layer, and Table A1 shows the matched record of longest and shortest axis length.



Figure 3:-. (a) Picture of microscope "Anyty V500IR/UV"-(a), (b) the measured longest axes lengths and shorteast axes length of particles (b), and (c) the reference ruler scale (c).

260 Snow density was measured using three instruments: snow tube, snow shovel and Snow-snow Fork 261 fork (Figure 4). The snow tube instrument, designed by Chinese Meteorological administration, contains 262 a metal tube with the base area of 100 cm^2 and the length of 60 cm, and a balance (figure 4a). It was 263 utilized to measure the snow density of a whole snowpack by weighing the snow sample. The snow 264 shovel is a 1500 cm³ wedge-type sampler, and its length, width and height are 20 cm, 15 cm, and 10 cm, 265 respectively (figure 4b). It was utilized to measure snow density every 10 cm (0-10 cm, 10-20 cm, 20-266 30 cm...). The Snow-snow Fork (figure 4c) is a microwave resonator that measures the complex 267 dielectric constant of snow, and adopts a semi-empirical equation to estimate snow density and liquid 268 water content based on the complex dielectric. The Snow ForkIt (figure 4e) was utilized to measure snow 269 density and liquid water content at 5-cm intervals starting 5 cm above the snow/soil interface (5cm, 10cm, 270 15 cm, 20cm...). In order to decrease the observation error, every measurement was conductedrepeated 271 three times. If there is was an abnormal value, the a fourth measurement would be performed to make 272 sureensure the accuracy. Table A2 is an example record table for snow density. The average value of the 273 three-time observation was the final value.



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Figure 4:-_Three instruments for snow density: (a) Snow tube (a), (b) Snow shovel (b), and (c) Snow Fork <u>fork(e)</u>.

280 2.2.3 Automatic radiation and temperature measurement

281 In the red field, the 4-component radiation was automatically measured by Component Net 282 Radiometer (NR01) manufactured by Hukseflux, and layered snow temperatures was measured by 283 Campbell 109S temperature sensors, respectively. The temperature sensors were set up on a vertical pole

whichpole __was vertically inserted in the soil (Figure 5). The heights of the sensors' heights are 0 cm, 5
 cm, 10 cm, 15 cm, 25 cm, 35 cm, 45 cm, and 55 cm above soil/snow interface, and temperatures were
 <u>collected-data</u>. The snow temperatures at these heights were collected every ten minute.

The NR01 net radiometer was set up to measure the energy balance between incoming short-wave and long-wave far infrared radiation versus surface-reflected short-wave and outgoing long-wave radiation. The range of short wave is 285~3000nm, and the range of long wave is 4.5~40_um. The 4component radiation was automatically recorded every ten minutes. In addition, the sensor is-was equipped with a Pt100 temperature sensor for parallel recording of the sensor temperature.



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Figure 5+. Set up of temperature sensors and CNR01 in the red field.

295 2.2.4 Meteorological observation

The meteorological data include air temperature, air pressure and humidity, wind speed, soil temperature at -5_cm, -10 cm, -15_cm and -20 cm₁ and soil moisture at -10 cm and -20 cm. These parameters are routine observations conducted at ANRMS, and were obtained through request from ANRMS. The instruments used for soil and weather parameters observations are produced by China Huayun Meteorological Technology Group <u>corporationCorporation</u>. The measurement parameters and their measurement instruments are listed in table 4.

302 Table 4. Automatically observed variables and the observation instruments, observation time and303 frequencies.

Observation
e time or
frequency
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5 Ten minute
5 Ten-minute
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1 component	Component Net	0.001 W/m ²	6 facts above		
4-component	Radiometer NR01		o leets above	Ten-minute	
radiation (w/m ⁻)	(Hukseflux)		ground		
Soil temperature	Soil temperature sensor	0.1 °C	-5cm, -10 cm, -	Houst	
(°C)	(China Huayun)		15cm and -20 cm	Hourry	
$\mathbf{S}_{\text{oil}} = \mathbf{S}_{\text{oil}} \mathbf{S}_{\text{oil}$	Soil moisture sensor	0.1%	-10 cm and -20	Houst	
Soli moisture (%)	(DZN3, China Huayun)		cm	Hourry	
Air temperature	Thermometer screen	0.1 °C	6 feet above	Handa	
(°C)	(China Huayun)		ground	Houriy	
Air mroggung (hDg)	Thermometer screen	0.1 hPa	6 feet above	Houst	
Air pressure (iiPa)	(China Huayun)		ground	Hourry	
A : h: dite (0/)	Thermometer	1%	6 feet above	Handa	
Air numidity (%)	screen(China Huayun)		ground	Hourry	
	Wind sensor(China	0.1m/s	10 m above	Houst	
wind speed (III/s)	Huayun)		ground	поипу	

The air temperature, pressure and humidity were collected using temperature and wetness sensor in 306 thermometer screen., Tthe wind speed and direction were measured using wind sensor set up at 10 m on 307 a tower. Soil moisture and temperature were automatically measured using moisture sensor and 308 temperature sensor. Figure 6 depicts the instruments for these observations.



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311 Figure 6:-_Instruments for observation of (a) air temperature and wetness (a), (b) wind speed (b), (c) soil 312 temperature (c) and (d) soil moisture (d).

313 **3 Description of consolidated IMCS dataset**

314 The microwave brightness temperature, snow parameters, meteorological data were recorded in 315 different formats, and their observation frequencies and times were different. These data must be 316 reorganized and consolidated for ease of use. The values from the three-time measurements for snow 317 density in each layer were averaged to obtain the final snow density. The length of the longest and shortest 318 axes of particles in each photo were measured using the software. The average lengths of longest and 319 shortest axes from all photos in each layer were obtained as the final grain size. The daily snow pit data 320 were finally consolidated into a NetCDF file "snow pit data.nc".

321 The time series of automated layered snow temperature and 4-component radiation data were firstly 322 processed with by removal of abnormal values and gap fill, and then were consolidated into a NetCDF 323 file "ten-minute 4 component radiation and snow temperature.nc". The ground-based brightness 324 temperatures and the formatted weather and soil data requested from ANRMS were provided 'as is'. 325 Brightness temperature data were divided into time series of brightness temperature and multi-angle 326 brightness temperatures, and separately stored in two NetCDF files, and tThe weather and soil data were 327 consolidated into a NetCDF file "hourly meteorological and soil data.nc". Table 3 describes the contents 328 of the provided dataset.

329 1) Brightness temperatures data:

Hinutely brightness temperature at 1.4 GHz, 18 GHz and 36 GHz for both polarizations at
 incidence angle of 50°. This data include date, time, incidence angle, azimuth angle, and brightness
 temperatures at the three bands for both polarizations.

2-Seventeen groups of calibrated brightness temperature at 1.4 GHz, 18.7 GHz and 36.5 GHz for
 both polarizations at different incidence angles (30, 35, 40, 45, 50, 55, 60°). This data include date<u>time</u>,
 incidence angles, azimuth angle, <u>and</u> brightness temperatures at the three bands for both polarizations.

336 2) Manual snow pit data:

Daily snow pit data include date, snow depth, layered snow thickness, average longest axis, average
shortest axis, grain shapes of each layer; layered snow density using snow fork (snow density at different
heights, such as SF_5cm, SF_10cm, SF_15cm), snow density using snow tube, layered snow density
using snow shovel (such as SS_0-10cm, SS_10-20cm, SS_20-30cm, SS_30-40cm).

- 341 3) Automated snow temperature and radiation data
- Ten-minute 4-component radiation and snow temperature data include date, time, short-wave incident radiation, short-wave reflected radiation, long-wave infrared incident radiation, long-wave infrared reflected radiation, sensor temperature, and snow temperatures at different heights (such as ST 0cm, ST 5cm)

346 4) Meteorological and soil data:

Hourly weather data include <u>date, hourtime</u>, air temperature, pressure, humidity, wind speed, soil temperature at 5 cm, 10 cm, 15 cm and 20 cm, and soil moisture at 10 cm and 20 cm.

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350 Table 3. Description of consolidated data

Data	Content	File name	Variables	
	Drichtmann		Time (yyyy-mm-dd hh:mm:ssYear, month, day, hour,	
Brightness	Brightness	TBdata.nc	minute, second), Tb1h, Tb1v, Tb18h, Tb18v, Tb36h, Tb36v,	
	temperature		incidence angle, azimuth angle	
temperature	Multi-angle		<u>Time (yyyy-mm-dd hh:mm:ss)</u> Year, month, day, hour,	
	brightness	TBdata-multiangle.nc	minute, second, Tb1h, Tb1v, Tb18h, Tb18v, Tb36h, Tb36v,	
	temperatures		incidence angle, azimuth angle	
			Time (yyyy-mm-dd Year, month, day), snow depth, th1, Lg1,	
	Layer thickness,		Sg1, th2, Lg2, Sg2, th3, Lg3, Sg3, th4, Lg4, Sg4, th5, Lg5,	
Manual snow	layered grain	Deile en en eit dete ne	Sg5, th6, Lg6, Sg6, Stube, SS_0-10, SS_10-20, SS_20-30,	
pit data	size and shape,	Daily snow pit data.nc	SS_30-40, SS_40-50, SF_5, SF_10, SF_15, SF_20, SF_25,	
	snow density		SF_30, SF_35, SF_40, SF_45, SF_50, shape1, shape2,	
			shape3, shape4, shape5, shape5	

Automated	1. common ant	Ten-minute 4	Time (yyyy-mm-dd hh:mm)Year, month, day, hour, minute,
snow	4-component	component radiation	SR_DOWN, SR_UP, LR_DOWN, LR_UP, T_Sensor,
temperature and	radiation, snow	and snow	ST_0cm, ST_5cm, ST_15cm, ST_25cm, ST_35cm,
radiation data		temperature.nc	ST_45cm, ST_55cm
	meteorological		Time (your me dd bb) Yeen menth day heur Tair Wein
Meteorological	data, soil	Hourly meteorological	<u>Time (yyyy-min-ut min) tear, month, uty, nour</u> , Tan, wan,
and soil data	nd soil data moisture and a	and soil data.nc	Pair, win, SM_10cm, SM_20cm, 1son_5cm, 1son_10cm, T_115120
	temperature		1 soii_15 cm, 1 soii_20cm

351 Note: th: snow thickness, Lg: long axis, Sg: short axis, shape: grain shape;

352 Stube: snow density observed using snow tube, SS: snow density observed using snow shovel, SF: snow density

353 observed using snow fork; ST: snow temperature; SR DOWN: downward short-wave radiation, SR UP: upward

354 short-wave radiation, LR DOWN, downward long-wave radiation, LR UP: upward long-wave radiation, T sensor:

sensor temperature; Tair: air temperature, Wair: air wetness, Pair: air pressure, Win: wind speed.

356 4 Overview and preliminary analysis of collected data from IMCS

357 4.1 Snow characteristics

358 4.1.1 Layering grain size and grain shape

During 2015/2016, snow cover began on 25-November 25, of 2015, and ended on March 25, of 2016. During this snow season, there were seven snowfall events and each formed a distinct snow layer except for the third event whose layering became was indistinguishable from the second layer (Figure 7 gray). The fourth event was the biggest, After, after which timethe snow depth started to decrease and the snow density increased. Snow cover began melting on March 14 and snow depth declined to zerodisappearedended within 10 days.

365 Grain sizes within all layers increased during the snow season, except in the bottom layer where 366 grain size experienced a decrease from December 28 to January 20 (Figure 8). In the vertical profile, 367 grain size increased from top to bottom with the snow age. The grain size of the fresh snow was 368 approximately 0.3 mm during the experiment. The biggest long and short axis, occurring in Layer 1-in 369 during the melt period, were up to 6 cm and 4 cm, respectively, and occurred in Layer 1 in during the 370 melt period. The length of short axes is approximately 0.7 of the length of long axes. The grain shape 371 generally developed from rounded grains to facet crystals, and then to depth hoar. After March 13, 2016, 372 the minimum air temperature increased to above 0°C, snowpack meltsnowmelt accelerated, and the grain 373 shape developed to melted forms (Figure 7).



Figure 7: Daily variation in snow layers and grain shape in each layer from November 27, 2015 to March 25, 2016.



Figure 8:-._Daily variation in grain size within each layer from November 27, 2015 to March 25, 2016. The thickness of each layer layer heights thicknesses are is presented in figure 97.

4.1.2 Snow density 383 Snow densityies measured by three different instruments shows that the density of fresh snow 384 ranged between 0.05~1.0 g/cm³ (Figure 9). The snow densities increased with snow age, and remained 385 stable after reaching ~0.2-0.25 g/cm3. From March 14 on, snow densities abruptly increased, and with 386 the maximum value reached was over 0.45 g/cm³. In the vertical profile, snow density increased from 387 top to bottom in the accumulation phase. However,, but after January 3, 2016, snow densities in the 388 middle layers were larger than those in the bottom and upper layers due to the well-developed depth hoar 389 of bottom layer. In the melting phase, there were no significant different for the snow densities in all 390 layers showed little difference. Snow fork provided most detail snow density profile, but it systematically 391 underestimated snow density compared with snow tube and snow shovel by 24% (Dai et al., 2022).



0.45 Snow density_snow shovel (g/cm³) snow shovel 0-10 cm snow shovel_10-20 cm snow shovel 20-30 cm 0.4 snow shovel_30-40 cm -snow shovel_40-50 cm 0.35 0.3 0.25 0.2 0.15 0.1 0.05 0 27-Nov 4-Dec 11-Dec 25-Dec 1-Jan 8-Jan 22-Jan 5-Feb 12-Feb 19-Feb 26-Feb -4-Mar 11-Mar 18-Mar (c) 25-Mar 18-Dec 15-Jan 29-Jan Date



November 27, 2015 to March 25, 2016₂, (a) overall snow density measured using snow tube, (b) snow density

394 395 396 397

398 399

4.1.3 Snow temperature

400 <u>The diurnal range of snow temperature decreased from top to bottom layers. As-and as the snow</u> 401 <u>depth increased, there were more layers with small diurnal variations (Figure 10).</u> Snow temperature at 402 0 cm (snow/soil interface temperature) showed <u>little-no significant</u> diurnal variation, remaining at 403 approximately -2.0 to 0.7_°C. Snow temperature in the top layer had the largest diurnal variation. The 404 diurnal range decreased from top to bottom layers and as the snow depth increased there were more 405 layers with small diurnal variations (Figure 10). After March 17, 2016, the snow temperature of all layers 406 were over 0°C, which meansimplying that snow cover did not refreeze anymore.

at 10-cm interval using snow shovel, and (c) snow density at 5-cm interval using snow fork.



Figure 10÷. Minutely variation in layered snow temperatures at 0 cm (snow/soil interface), 5 cm, 15 cm, 25
cm, 35 cm, 45 cm and 55 cm above ground during experiment time.period

410 **4.2 Soil temperature and moisture**

The soil temperature at 5 and 10 cm remained stable and below 0 °C during the snow season but presented large fluctuation before (after) snow on (off) (Figure 11). The temperature difference between 5 cm and 10 cm was much larger before snow cover onset than during snow cover period. The soil moistures at 10 cm were above 10% before snow cover onset and after snow off., and tThere were two soil moisture peaks within the snow cover period, one from December 12-14 and another the other from January 1- 20, within the snow cover period.



417

Figure 11÷. (a) Hourly soil temperature at 5 cm, 10 cm, 15 cm and 20 cm below the snow/soil interface (a), and (b) soil moisture at 10 cm and 20 cm below the snow/soil interface (b).

420 4.3 Brightness temperature

421 The microwave brightness temperatures varied with snow-and, soil-characteristics, and weather 422 conditions. Figure 12 shows the daily brightness temperatures, brightness temperature difference 423 between 18 and 36 GHz, and snow depth at 1:00 am local time. Figure 13 shows the hourly variation in 424 brightness temperatures at 1.4 GHz, 18.7 GHz and 36.5 GHz and air temperature after February 1. Data 425 show that Tb36h and Tb36v decreased during the full snow season, Tb18h shows-exhibited an obvious 426 decline after February 18, and Tb18v after March 3 (Figure 12). After January 4, though snow depth 427 stopped increasing, but the brightness temperature continued to decrease and brightness temperature 428 difference increased. Based on Figure 8, snow density became stable on-after January 15. Therefore, after 429 January 4, the decreasing brightness temperatures was mainly caused by growing grain size. The 430 variation of L band mainly relies on soil moisture and soil temperature change. We have soil temperatures

431 <u>at 0 cm, 5 cm and 10 cm and soil moisture at 10 cm. However, the L band reflects the soil moisture within</u>
 432 <u>5 cm which was absent in this experiment.</u>

After February 25, brightness temperature exhibited a distinct cycle of daytime increase and nighttime decrease (Figure 13), resulting from large liquid water content caused by high daytime air temperature (above 0_°C) and the melted snowpack refreezing at nighttime. After March 14, there was another big rise in air temperature and even the nighttime air temperatures were above 0°C. During this period of accelerated snowmelt, <u>as the liquid water within the snowpack did not refreeze completely at</u> night<u>a</u> and both the brightness temperatures <u>at three bands</u> and brightness temperature difference exhibited irregular behavior.

440 The variation of L band <u>mainly rely on</u> was mainly influenced by soil moisture and soil temperature 441 <u>change</u>. We have soil temperatures at 0 cm, 5 cm and 10 cm and soil moisture at 10 cm. However, the L 442 band reflects the soil moisture within 5 cm which was absent in this experiment. Actually, we did not 443 find the variation of brightness temperature at L band had relationship with soil moisture at 10 cm and 444 soil temperature.



Figure 12÷. Daily variations in brightness temperatures at 1.4 GHz, 18 GHz and 36 GHz, for horizontal
(Tb1h, Tb18h, Tb36h) and vertical polarizations (Tb1v, Tb18v, Tb36v), and the differences between Tb18h
and Tb36h (Tb18h - Tb36h), and between Tb18v and Tb36v (Tb18v - Tb36v), at 1:00 am (local time), from
November 27, 2015 to March 26, 2016. (a) for horizontal polarization, and (b) for vertical polarization.



456



Figure 13. Hourly variation in (a) Tb1h, Tb18h, Tb36h, Tb1v, Tb18v, and Tb36v (a), and (b) air
temperature, soil moisture at 10 cm and soil temperature at 5 cm, and daily variation in snow depth (b),
from February 1 to March 28, 2016.

457 The brightness temperatures at 18.6-7 and 36.5 GHz from AMSR-2 and at 1.4 GHz from SMAP 458 were _-compared with the ground-based observations at the overpass time (Figure 14). Although there 459 were large differences between satellite and ground-based observations, the general temporal patterns 460 are were the same, same. Specifically, even the abrupt changes at 18.7 and 36.5 GHz between on March 461 3 and March 4-16 is were captured by both satellites and ground-based sensors. Brightness temperatures 462 at 1.4 GHz from both SMAP and ground microwave radiometer kept stable before March 16, after when, 463 brightness temperature rapidly decreased because of the increase of liquid water contentsoil moisture, 464 The correlation coefficients at both polarizations were approximately 0.96, 0.7 and 0.88 for 36.5 GHz, 465 18.6-7 GHz and 1.4 GHz, respectively. Satellite observed brightness temperature presented less decrease 466 trend than ground-based observation., and tThe difference at 36.5 GHz is larger than those at 18.6-7 and 467 1.4 GHz, Brightness temperatures at 1.4 GHz from both SMAP and ground microwave radiometer kept 468 stable before March 16, after when, brightness temperature rapidly decreased because of the increase of 469 liquid water content. The difference between ground-based and satellite observation might be attributed 470 to the different different field of viewing areas.



472 473

Figure 14:-<u>.</u>Comparison of brightness temperature between ground-based<u>(0)</u> and satellite-based observation<u>s</u> (s: satellite; o: observation), (a) for 36 GHz, (b) for 18 GHz, <u>and</u>(c) for 1.4 GHz<u>.</u>

474 4.4 4-component Radiation

The land surface albedo is strongly related to the land cover. In this experiment, the downward short-wave radiation presented general increase after January, <u>and-while</u> the trend became distinctive after February (Figure 15). The upward short-wave radiation abruptly increased when the ground was covered by snow (after November 21), and sharply declined on the snow off day (March 25). From the

479 first snowfall by the end of January, the ratios between upward and downward short-wave radiation were
480 approximately 95%. The ratio decreased with snow age, and in the end of snow season, the ratios
481 decreased to <u>approximately below 5025% on snow off day because of increasing melted water</u>.



482

Figure 15+. Minutely variation in 4-component radiation and daily variation in snow depth at Altay station
from November 3, 2015 to April 15, 2016.

485 5 Discussion

486 5.1 Applications

Although the<u>Our</u> dataset is, though-just for one <u>snow</u> season-observation, the <u>provides</u> daily snow pit observations with coincident microwave and optical radiation data, <u>including comprehensive and in</u> a full snow season provide the most detailed variation of snow parameters, which allow<u>ing</u> researchers to find more details in snow characteristics and their relationship with remote sensing signatures. The dataset also fills the snow observation gap in mid-low snow depth area with relative short snow cover duration.

493 The snow pit data and microwave brightness temperatures have proven useful for evaluating and 494 updating a microwave emission transfer model of snowpack (Dai et al., 2022). This dataset reflected the 495 a general fact that brightness temperature at higher frequencies presented stronger volume scattering of 496 snow grains, and were more sensitive to snow characteristics. This experiment revealed that the dominant 497 control for the variation of brightness temperature was the variation of grain size but not the snow depth 498 or SWE. In the stable period, The largest snow depth or SWE did not correspond to the largest brightness 499 500 of grain size, the maximum difference occurred bwas increasing increased with growing grain size in the 501 condition of dry snowpackefore melting for stable snow cover. Therefore, the daily snow depth variations 502 curve derived from passive microwave remote sensing datasets tend to exhibit a temporal offset from 503 those of in situ observation.

504 During the snow season, brightness temperatures for both polarizations presented similar variation 505 <u>trendss</u>, <u>but theythough behaved behaving</u> different in <u>some time periodsfluctuation</u>. The horizontal 506 polarization was more sensitive to environment and was less stable than vertical polarization. Besides, 507 the polarization difference at 18.7 GHz and 36.5 GHz showed increase and decrease trends, respectively, 508 during the experimental period. <u>The This phenomenon was different from results for 18 GHz were</u> 509 <u>opposite to the simulation results (Dai et al., 2022)</u>. The different polarization behavior at 18.7 and 36.5 510 GHz might be related to the environmental conditions, snow characteristics and soil conditions. However, 511 <u>as</u> the subsurface soil moisture was not observed, the dynamic ground emissivity could not be estimated. 512 <u>As it is known, as</u> L band has strong penetrability, and the brightness temperature variations were 513 predominantlyly related to subsurface soil conditions, except whenunless for the situations that the liquid 514 water content within snowpack was high. Therefore, in the condition of soil moisture data absence, L 515 band brightness temperatures were expected to reflect soil moisture variation which that influences the 516 soil transmissivity (Babaeian et al., 2019; Naderpour et al., 2017; Hirahara et al., 2020).

517 Snow surface albedo significantly influences the incoming solar radiation, playing an important role 518 in the climate system. The factors altering snow surface albedo contains the snow characteristics (grain 519 size, SWE, liquid water content, impurities, surface temperature etc), external atmospheric condition and 520 solar zenith angle (Aoki et al., 2003). Snow albedo was estimated based on snow surface temperatures 521 in some models (Roesch et al., 1999), while others considered that snow surface albedo was mainly 522 related to to depend mainly on snow aging age (Mabuchi et al., 1997). In this experiment, we obtained 523 the 4-component radiation, snow pit and meteorological data,- These data-provideing nearly all 524 observations of possible influence factors, and therefore could be utilized to discuss and analyze 525 shortwave radiation process of snowpack, and validate or improve multiple-snow-layer albedo models.

526 Snow grain sizes and snow densities within different layers presented different growth rates during 527 different time periodsperiods. Generally, the growth rates are related to the air temperature, pressure and 528 snow depth (Chen et al., 2020; Essery, 2015; Vionnet et al., 2012; Lehning et al., 2002); therefore, this 529 dataset can be used to analyze the evolution process of snow characteristics, as well as validation data 530 for snow models.

5.2 Uncertainties

531

532 During the experiment, some uncertainties were produced due to irresistible factors. It is-was 533 reported that the sampling depth of the L-band microwave emission under frozen and thawed soil 534 conditions is determined at 2.5 cm (Zheng et al., 2019). We did not collect subsurface soil moisture, and 535 the L band radiometer observation began on January 30, 2016. Therefore, it is difficult to obtain the 536 ground emissivity in the full snow season based on the data. The soil moisture data at 10 and 20 cm under 537 soil/snow interface cannot be directly used to validate and develop soil moisture retrieval from L band 538 brightness temperature. We hope In the future, detailed soil moisture profile will be observed to estimate 539 the subsurface soil moisture to fill the gap.

The grain size data were collected through taking photos. When measuring the length of grains, the grain selection has subjectivity, and the released data are <u>average valuesstatistic results based on the</u> ehosenrecorded grain sizes. Although the general variation trend can be reflected by the time series of average grain size, some details might be missed. Therefore, <u>for who with interests</u>, the original grain photos could be provided through requesting for authors. In snow-melt period, large liquid water content would influence the measurement results of snow fork. <u>SoTherefore</u>, it is suggested to use small-size snow shovel or cutter to observe layered snow density in future experiments.

547 One season observation is quite valuable for developing and validate remote sensing <u>snow retrieval</u> 548 method_–or snow model, although the representativeness of this observation <u>remains unknownrequires</u> 549 <u>further analysis</u>. <u>Nevertheless, We need</u> more years of observations <u>should be considered to increase the</u> 550 <u>statistical significance of to endorse or confirm</u> the evolution of snow characteristics.

551 6 Conclusions

In a summary, the IMCS campaign provides a time series of snow pits observation, meteorological parameters, optical radiation and passive microwave brightness temperatures in <u>the the</u> snow season of 2015/2016. The dataset is unique in providing microwave brightness temperatures and coincident daily snow pits data over a full snow season at a fix site. <u>The first use of our dataset is for the validation of</u> <u>snow microwave emission models, improving its simulation accuracy.</u>

557 The daily snow pit data which provide a detail description of snow grain size, grain shape, snow 558 density and snow temperature profiles. Generally, grain size grew with snow age, and increased from top 559 to bottom. Snow grains are rounded shape with small grain size in the top layer, and depth hoar with 560 large grain size in the bottom layer. Snow density experienced increase-stable-increase variation, and the 561 densities of the middle layers were greater than the bottom layer due to the well-developed depth hoar in 562 the stable period. The data can be used to analyzes the evolution process of snow characteristics 563 combining with weather data, also for the validatione and improvement of the snow process models, such 564 as SNOWPACK (Lehning et al., 2002), SNTHERM (Chen et al., 2020), etc. The improvement of these 565 snow process models can further enhance the prediction accuracy of land surface process and hydrology 566 models, and the simulation accuracy of snow microwave emission models.

567 Microwave radiometer data and snow pit data have been utilized to analyze the volume scattering 568 features of snow-pack at different frequencies (Dai et al., 2022). Results showed that grain size is-was 569 the most important factor to influence snow volume scattering. The data can also be used to <u>for analysis</u> 570 <u>of further analyze</u>-polarization characteristics of snow-pack, combining with soil and weather data, and 571 be used to validate different microwave emission models of snowpack.

572 The microwave and optical radiations were simultaneously observed. Existing studies reported that 573 the optical equivalent diameter must be used in microwave emission model with caution (<u>LöweLowe</u> 574 and Picard, 2015; Roy et al., 2013). These data provide a <u>good new</u> opportunity to analyze the difference 575 between the influence of grain size on microwave and optical radiation, establishing the bridge between 576 effective optical grain size and microwave grain size.

577 7 Data availability

The IMCS consolidated datasets are available after registration on the National Tibetan Plateau Data Center and available online at http://data.tpdc.ac.cn/zh-hans/data/df1b5edb-daf7-421f-b326cdb278547eb5/ (doi: 10.11888/Snow.tpdc.270886). Microwave radiometry raw Data are available for scientific use on request from Northwest Institute of Eco-Environment and Resources, Chinese Academy of Sciences.

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585 Author contributions: LD and TC designed the experiment. LD, YZ, JT, MA, LX, SZ, YY YH and LX 586 collected the passive microwave and snow pit data. HL provided the 4-component radiation and snow 587 temperature data. LW provided meteorological data. LD write the manuscript, and TC made revision. All 588 authors contributed to the data consolidation.

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590 **Competing interests:** The authors declare that they have no conflict of interest.

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Appendix



Strati	Thickn					_					
graph	ess	Shape		Grain size (longest axis * shortest axis)(mm)							
у	(cm)										
			0.595	0.472	0.450	0.615	0.374	0.647	0.656	0.544	0.717
			*0.43	*0.47	*0.43	*0.47	*0.31	*0.30	*0.52	*0.51	*0.44
the			6	1	6	4	4	7	9	9	7
fifth	3cm	#22	0.750	1.056	0.623	0.451	1.397	1.235	0.600	0.633	0.729
			*0.44	*0.95	*0.37	*0.40	*0.63	*0.32	*0.42	*0.55	*0.42
			5	5	8	5	5	7	1	6	3
			2.605	1.850	1.626	1.767	1.718	2.255	1.674	1.542	3.505
			*2.01	*1.32	*1.55	*1.68	*1.53	*1.29	*1.60	*1.26	*1.44
			1	8	4	5	5	6	1	9	0
the	3cm	#37	3.055	1.440	2.461	2.757	2.179	2.393			
fourth			*1.77	1.448	*1.91	*2.11	*2.05	*1.78			
			4	*1.37	4	5	9	8			
			2.569	2.073	2.591	1.869	2.067	1.209	1.719	1.648	1.911
			*1.60	*2.13	*1.41	*1.80	*1.26	*1.10	*1.18	*0.97	*1.58
			7	0	4	2	6	6	8	5	2
		#27,	1.921	1.518	1.291	1.690	1.756	1.812	1.733	1.880	2.411
the	25cm	#31,	*1.71	*1.06	*1.14	*1.55	*1.39	*1.26	*1.67	*1.51	*1.22
third		#37	0	7	7	1	8	3	2	8	0
			2.118	1.614	1.795	2.215	1.864	1.967	2.008	1.362	1.484
			*1.72	*1.45	*1.70	*2.31	*1.69	*1.65	*1.39	*1.14	*1.29
			7	7	5	1	2	1	5	1	1
			4.251	2.012	2.805	1.799	1.402	3.040	2.850		
			*2.26	3.012	*1.99	*1.41	*1.19	*2.07	*2.09		
the	10	#33,	6	*2.65	5	5	5	3	5		
secon	12	#34	3.900	2.420	2.515	2.044	2.506	2.894	2.413	2.494	4.929
d			*2.53	*2.33	*2.20	*2.03	*2.36	*2.16	*1.95	*1.81	*3.25
			2	3	6	2	3	1	0	6	7
			4.933	3.207	3.562	2.818	3.581	6.179			
the	4	#40,	*3.37	*2.77	*1.70	*1.66	*2.51	*3.56			
first	4	#34,	8	4	1	8	8	2			
		#38									

C10	T_{-} 1 1 A 1 D 1 1 1 4 1 4 4 4 4 4 4 4 4
n 1 /	I able A L. Recorded longest and shortest axis length in Figure A.
016	The second and solid of the second se

620 Table A2: <u>One An</u> example of record table for snow density observation.

621

observation date:	20160111	observation time: 9:03-9	:40	weather: clear	snow depth: 48cm	
	Snow Folk table			Snow tube t	able	
observation height (cm)	liquid water content(%)	snow density (g/cm3)	snow depth(cm)	46.5	47	47.5
	0	0.1923	snow pressure(g/cm2)	9.1	9	9.5
5	0.118	0.1882	snow density(g/cm3)	0.1957	0.1915	0.2000
	0	0.1882				
	0.461	0.164		snow shovel	table	
10	0.46	0.1631	observation layer (cm)	weight of shovel+snow(g)	weight of shovel(g)	snow density(g/cm3)
	0.461	0.1361		865.04	572.16	0.1953
	0.123	0.2532	0-10	858.72	572.16	0.1910
15	0	0.2506		866.69	572.16	0.1964
	0	0.2417		878.58	572.16	0.2043
	0.24	0.2159	10-20	887.04	572.16	0.2099
20	0.119	0.2155		872.79	572.16	0.2004
	0.119	0.2146	20-30	905.34	572.16	0.2221
	0.117	0.1977		903.41	572.16	0.2208
25	0	0.1994		907.88	572.16	0.2238
	0	0.1984		832.75	572.16	0.1737
	0	0.1919	30-40	838.14	572.16	0.1773
30	0	0.1966		837.27	572.16	0.1767
	0	0.1928				
	0	0.1534	40-50			
35	0	0.1517				
	0	0.1472				
	0.325	0.1097	50-60			
40	0	0.1054				
	0.107	0.1088				
	0	0.0922				
45	0	0.0991				
	0	0.0928				
50						

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