Attenuated atmospheric backscatter profiles measured by the CO_2 Sounder lidar in the 2017 ASCENDS/ABoVE airborne campaign

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Abstract. A series of attenuated atmospheric backscatter profiles were measured at 1572 nm by the CO_2 Sounder lidar during the eight flights of the 2017 ASCENDS/ABoVE (Active Sensing of CO_2 Emission over Nights, Days, & Seasons, Arctic Boreal Vulnerability Experiment) airborne campaign. In addition
10 to measuring the column average CO_2 mixing ratio from the laser signals reflected by the ground, the CO_2 Sounder lidar also recorded the height-resolved attenuated atmospheric backscatter profiles beneath the aircraft. We have recently processed these vertical profiles with a 15-m vertical spacing and 1-s integration time along the flight path (~200 m) for all the 2017 flights and have posted the results at NASA Distributed Active Archive Center (DAAC) for Biogeochemical Dynamics (Sun et al., 2022),
15 https://doi.org/10.3334/ORNLDAAC/2051. This paper describes the measurement principles, the data processing technique, and the signal to noise ratios.

1 Introduction

NASA Goddard Space Flight Center (GSFC) developed the pulsed multi-wavelength CO_2 Sounder lidar for measuring column average CO_2 mixing ratio (XCO_2) near 1572 nm (Abshire et al., 2018) as an
20 airborne demonstrator for NASA's planned Active Sensing of CO_2 Emission over Nights, Days, & Seasons (ASCENDS) space mission (Kawa et al., 2018). The airborne CO_2 Sounder lidar has operated successfully during several campaigns, including in the 2017 Arctic Boreal Vulnerability Experiment (ABoVE) airborne campaign onboard the NASA DC-8 aircraft.

The CO_2 Sounder lidar uses a pulsed laser and rapidly step-tunes the laser wavelength across the
25 1572.335 nm CO_2 absorption line. The lidar receiver uses a single photon sensitive HgCdTe avalanche photodiode (APD) (Sun et al., 2017) and an analog-to-digital convertor (ADC) which records the entire atmospheric backscatter profile as well as the surface reflected signal. The surface returns are used to retrieve the XCO_2 of the atmosphere column travelled by the laser pulse (Sun et al, 2021).

Recently, we have processed the airborne atmospheric backscatter measurements to provide the
30 atmospheric backscatter profiles from the airplane flight altitude to the surface with a 15-m vertical sampling interval for all eight flights of the 2017 campaign. The results have been posted on the repository for NASA Distributed Active Archive Center (DAAC) for Biogeochemical Dynamics (Sun et al., 2022). Although the backscatter profiles are not used directly in the XCO_2 retrieval, they provide important context for interpreting the retrieved XCO_2 measurements. They can be used to identify

35   clouds for retrieving partial column XCO_2 to cloud tops (Mao et al., 2018). They also enable identifying and profiling smoke plumes from wildfires to help assess their impact on XCO_2 (Mao et al. 2021).

        This paper describes the data processing of the airborne CO_2 Sounder lidar's atmospheric backscatter profiles from the eight flights of the 2017 airborne campaign. We first give a brief description of the lidar and its data structure. We then describe the details of the Level-0 and Level-1

40   data processing and the signal to noise ratio (SNR) of the profiles.

2. A brief description of the CO_2 Sounder lidar instrument

2.1 The airborne CO_2 Sounder lidar

The airborne CO_2 Sounder lidar transmits laser pulses toward nadir from the aircraft and detects and records the laser signals backscattered from the atmosphere and the surface. The wavelength of the

45   single frequency laser is rapidly step-tuned across the CO_2 absorption line centered at 1572.335 nm. The transmitted laser pulse energies are also measured and are used to normalize the received signal to correct for variations in laser power at the different wavelengths. The lidar receiver digitizes the pulse waveforms from atmospheric backscatter as well as the surface reflection. Figure 1 shows a simplified instrument block diagram. More details about the airborne CO_2 Sounder lidar can be found in Abshire et

50   al. (2018).

        The laser emits 1-µs wide pulses with a nearly rectangular pulse shape continuously at 10 kHz. The laser wavelength is step-tuned across the CO_2 absorption line for 30 pulses, followed by 2 pulses during the period of wavelength rewind. The wavelengths of the 30 laser pulses during a wavelength scan are listed in Table 1. The laser wavelengths used by the lidar are offset-locked to the CO_2

55   absorption line in a gas cell at 40 mbar pressure and 296 K temperature. The absorption line center is at 1572.335 nm according to HITRAN 2012 (Numata et al., 2011, 2012). The offset locking frequency is changed between pulses to step the laser wavelength across the CO_2 absorption line.

        Figure 2 shows the laser wavelengths plotted across a CO_2 absorption line shape that was computed from a modelled atmosphere. Note that the distribution of laser wavelengths is slightly offset

60   from the line center of the absorption line due to the difference in pressure between the CO_2 in the reference cell and that of the atmosphere being modelled and measured. The laser pulse emission times, the wavelength scan and the data acquisition are all synchronized to the Coordinated Universal Time (UTC) by the on-board computer. The lidar receiver employs a photon-sensitive HgCdTe APD detector (Sun et al., 2017). The detector has a linear analog response and has sufficient sensitivity and linear

65   dynamic range to record both the time-resolved atmospheric backscatter profiles and the surface reflected signals. The output of lidar detector is digitized continuously with 16-bit resolution at 100 MHz sampling rate.
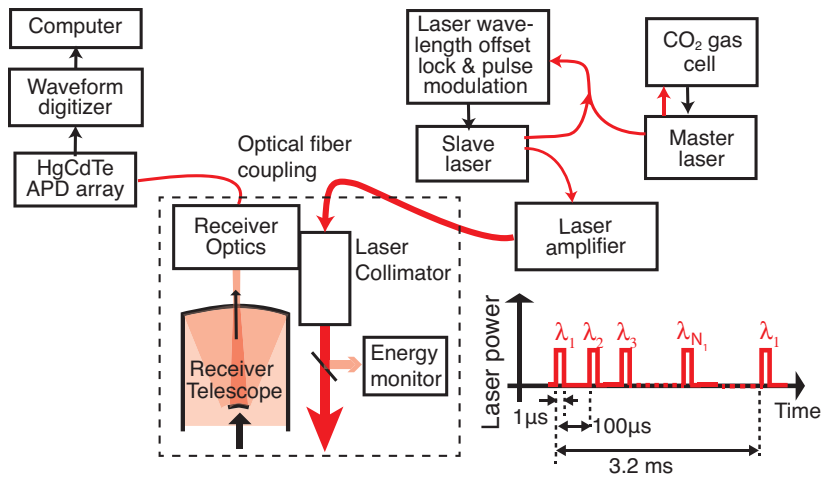
Figure 1. A simplified block diagram of the airborne CO_2 Sounder lidar.

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Table 1. List of laser wavelengths used in the 2017 airborne campaign for the CO_2 Sounder lidar. All wavelengths are given as the difference of the actual wavelength from the center of the 1572.335 nm CO_2 absorption line in the reference cell.

Laser pulse number	Frequency and Wavelength difference from the line center of the CO_2 cell		Laser pulse number	Frequency and Wavelength difference from the line center of the CO_2 cell	
	GHz	Picometer		GHz	Picometer
1	12.25	-101.013	16	-0.25	2.062
2	9.50	-78.388	17	-0.50	4.123
3	8.25	-68.031	18	-0.75	6.185
4	6.25	-51.539	19	-1.00	8.247
5	4.25	-35.047	20	-1.25	10.308
6	3.25	-26.801	21	-1.50	12.370
7	2.50	-20.616	22	-1.75	14.432
8	2.00	-16.493	23	-2.00	16.493
9	1.75	-14.431	24	-2.50	20.617
10	1.50	-12.370	25	-3.25	26.802
11	1.25	-10.308	26	-4.25	35.048
12	1.00	-8.246	27	-6.25	51.542
13	0.75	-6.185	28	-8.25	68.037
14	0.50	-4.123	29	-9.50	78.346
15	0.25	-2.062	30	-12.25	101.026

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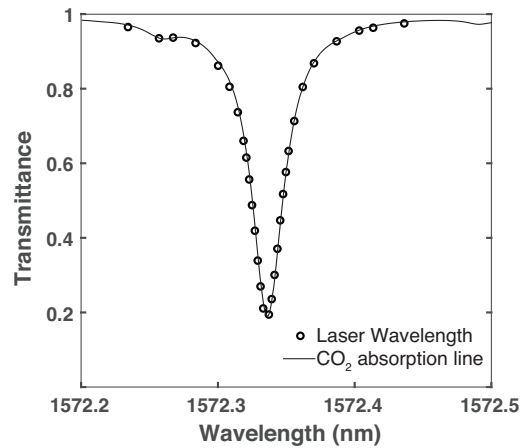

80 **Figure 2.** The laser wavelengths of the CO_2 Sounder lidar overlaid on a typical CO_2 absorption line computed for the nadir path using a modelled atmosphere.

2.2 Data structure of the airborne CO_2 Sounder lidar

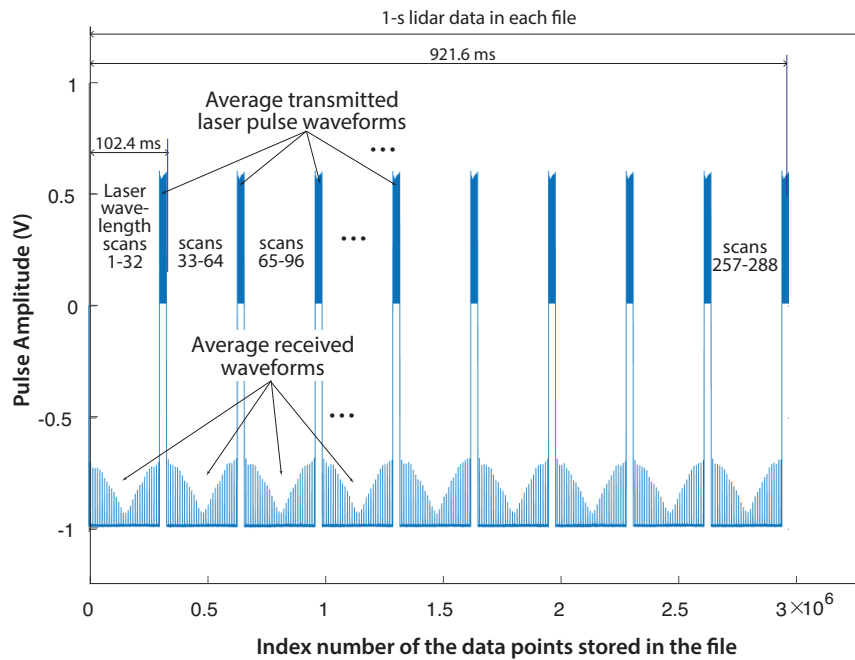
The CO_2 Sounder lidar operations are synchronized with the UTC time via a GPS receiver. The measured signals are stored into data files at the end of every UTC second. The aircraft housekeeping and the GPS data are also recorded every UTC second. Each lidar data file contains 9 groups of 30 received backscatter profiles, one profile for each wavelength averaged over 32 wavelength scans. Therefore, each group contains 102.4 ms of measurement data (~10 Hz sample rate) and each 1-s data file contains measurement data from 32×9 wavelength scans. A DC offset of -1.1 V is added to the received waveforms before the ADC to make use of its full dynamic range (± 1.25 V). At the 10 kHz pulse rate, the time between transmitted laser pulses is 100 μs. At 100 MHz ADC sampling rate, each waveform segment contains 10,000 ADC sample points. The transmitted laser pulse waveforms are sampled and averaged by another ADC at the same sampling rate for 1,000 points without adding a DC offset. The lidar controller suspends the laser wavelength scan after all the 9 groups of data (total 921.6 ms) are collected. The remaining time before the next whole UTC second is used to generate the timestamp and combine the transmitted laser pulse waveforms into the same file. The pulse waveforms from the 31st and 32nd laser pulses during the wavelength rewind are discarded. The transmitted pulse waveforms are appended to the 30th received waveform. Figure 3 shows a sample plot of recorded data file in the 16-bit signed integer format used by the ADCs.

F**igure 3.** Plot of the recorded data file structure for CO_2 Sounder lidar. The data were taken on 8/8/2017 23:30:00 (UTC) over the mountains as the aircraft flew toward the Edwards Airforce Base in California at an altitude of 11.2 km. There are 9 groups of 30 pulse waveforms, one for each laser wavelength in the scan. Each is averaged over 32 scans. The corresponding waveforms for the 30 transmitted laser pulses, which are sampled at the same rate for 1,000 points each, are appended to the end of the 30th received pulse waveform.

Figure 4 shows a plot of the first 30 individual pulse waveforms from the first group taken on 8/8/2017 at 23:34:00 (UTC). The first group of pulse waveforms with relatively low amplitude are from the scatter from the aircraft's nadir window. These are used as the time reference in the laser pulse time-of-flight measurements. The second group of waveforms about 50 μs after the window returns is from thin clouds, and the last set of waveforms are from the ground.

The laser pulse shape is nearly rectangular but its peak power gradually decreases over the 1 μs pulse width period. This is caused by the gradual depletion of energy stored in the fiber laser amplifier during the pulse. The detector has a slight ringing after the relatively strong ground return, which is omitted in the subsequent signal processing. There is also a small baseline offset from the detector in addition to the -1.1 V DC offset added before the ADC. The total baseline offset is estimated by averaging the waveform segment before the transmitter pulses and removed in data processing. The 30 pulse waveforms reflected by the window have nearly the same pulse amplitude. The cloud and ground return pulse waveforms show increasing variation in the detected pulse amplitude caused by atmospheric CO_2 absorption as the laser wavelengths step through the CO_2 absorption line.

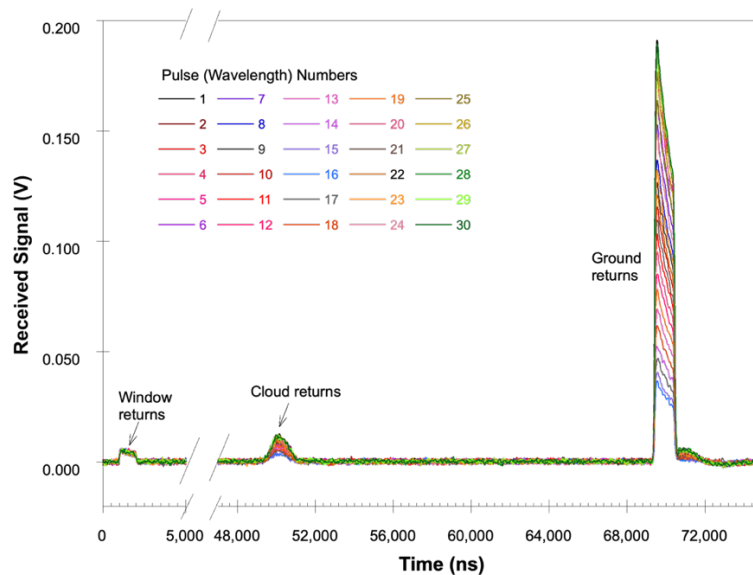



Figure 4. Overlay of the 30 received pulse waveforms recorded by the airborne lidar receiver on 8/8/2017 at 23:34:00 (UTC).

The XCO_2 is retrieved from the received laser pulse energies reflected by the surface. At each wavelength the received pulse energy is calculated by first removing the DC offset, then integrating the pulse waveforms reflected by the ground surface and normalizing them by transmitted pulse energies. The relative atmospheric transmissions for each of the laser wavelengths are calculated as the normalized received pulse energies divided by the square of the range measured by the lidar. The surface reflectance is assumed constant during the wavelength scan. Changes in the received pulse energies at different wavelengths are attributed to the absorption by atmospheric CO_2. The XCO_2 value is retrieved via an algorithm using a linear least-squares fit of the modeled CO_2 line shape to the lidar sampled absorption line shape. Details about the signal processing for the XCO_2 retrieval can be found in Sun et al. (2021).

The signals used to calculate the atmospheric backscatter profile are contained in the same data files and are obtained by processing the signal waveforms before the ground returns. Some preliminary backscatter profiles in raw ADC units (integers) were reported earlier for this campaign (Allan et al., 2018). We have since then calibrated the backscatter profiles and released results on NASA Distributed Active Archive Center (DAAC) for Biogeochemical Dynamics (Sun et al., 2022). Details of the data processing and the estimated SNR of the profiles are presented in the remainder of this paper.

3. Data processing of the atmospheric backscatter profiles from the CO_2 Sounder lidar

3.1 Level-0 Data Processing

The Level-0 data processing converts the raw ADC output in 16-bit integers to the atmospheric backscattered signal in the physical units (V) of the detector output. It also removes all the known

artifacts in the data such as baseline DC offsets. To improve the SNR, we averaged the backscatter
150 profiles measured for laser wavelengths 2, 3, 4, 27, 28, 29 and 30 (7 in total). We did not include the
first wavelength of the scans because occasionally the laser operation was not fully settled after the
wavelength rewind. As shown in Figure 2, these wavelengths undergo only small CO_2 absorption, so
this average may be considered as a backscatter measurement at "off-line" wavelengths.

For each received pulse waveform, the DC offset from the detector is first subtracted by
155 calculating the average detector signal for the time interval that occurs immediately before the return
from aircraft window. Since the pre-window-return region is long after the ground return from the
previous laser pulse but before the next one, it is composed primarily of detected solar background, the
dark noise from the detector, and the DC offset of the lidar receiver. One component of the DC offset is
a constant offset (-1.1 V) added to the detector output before the ADC. The other component is the
160 inherent DC offset from the detector, which can drift slowly over time. It also can change after the
detector is overloaded by excessively strong signals, such as returns from clouds immediately under the
aircraft or returns from specular reflection at the ground.

The transmitted pulse energies are calculated from the transmitted pulse waveforms by first
subtracting the DC offset and then integrating the signal over the pulse period. They are then used to
165 normalize the corresponding received waveform. The average transmitted pulse energy over the entire
flight is also recorded in the Level-0 data file. A known time delay within the lidar electrical and optical
system (corresponding to a range of 26.4 m) is removed next. Following this, received pulse waveforms
in the 1-s data file are averaged together to improve the SNR. Finally, the averaged pulse waveform is
converted into the units of the detector output (V) by dividing the waveform data by the scale factor of
170 the ADC.

Screening is then performed to eliminate any data with the following abnormalities: the
transmitted pulse waveforms are missing in the data file; the ground return pulse is saturated with pulse
amplitude above 1.1 V after removing the DC offset; or data are taken while the detector is still
recovering from saturation where the estimated detector DC offset falls below 0 V or above 0.5 V after
175 removing the known 1.1 V offset (nominal DC offset from the detector is 0.1 to 0.2 V). The mean and
standard deviation of the waveform within the pre-window region for each averaged waveform are then
recorded in the Level-0 data files in case further screening is needed. Lidar data that contain cloud
returns within 3-km of the aircraft window are also flagged because within this distance the laser beam
does not completely overlap the receiver field of view (Allan et al., 2018). The data within this range
180 needs to be further calibrated to account for the losses due to the overlap factor.

The aircraft's navigation data are then merged with the lidar data. The aircraft flight data are
obtained from archived airborne campaign housekeeping data gathered during the flight from the on-
board GPS receiver and other aircraft instruments. The following parameters are extracted from the
archived DC-8 housekeeping data for each second: the UTC time, latitude, longitude, altitude, pitch and
185 roll angles, and the DC-8's radar-measured range to the surface in nadir direction.

The effect of off-nadir pointing is corrected next. The lidar is mounted to the aircraft's deck and
the laser beam is pointed down perpendicular to the deck. The aircraft flies slightly pitched up during
cruise. It also performs pitch and roll maneuvers from time to time. When the laser beam is pointed
away from nadir, the laser pulse travels a slant path (a longer distance) before reaching the ground. This
190 effect, if not corrected, would result a longer length profile and a lower ground elevation when plotted

with other profiles taken along nadir direction. To correct this effect, the range bin size (15 m) is first multiplied by the cosine of the combined off-nadir pointing angle. This has the effect of compressing the profile and correcting the elevation of the ground return. The resultant is then resampled by interpolating at 15-m intervals in the nadir direction. The attenuated backscatter coefficients in the corrected range bins should also be corrected for the additional attenuation by the atmosphere above the layer. Under uniform layer structure, atmospheric transmission from the aircraft to the given altitude equals $T(\theta_{nadir}) = [T(0)]^{1/cos(\theta_{nadir})}$ with θ_{nadir} the laser beam pointing angle from nadir. The aircraft pitch angle was mostly within 5 degrees during cruise. The roll angle was near zero except during sharp turns or spiral-dows. Total decrease in the two-way transmission is < 5% under this condition. We did not correct this effect in the current data release. Since the combined off-nadir angle is relatively small, it is a relatively small effect for this data set. Additionally, since both the pitch and roll angles are included in the released data, the user may apply this correction if necessary.

The surface elevation is calculated as the difference between the aircraft altitude from the on-board GPS receiver and the range measured by the lidar after correcting for the effect of off-nadir pointing angle. Missing points in the surface elevation, such as those obscured by clouds, are interpolated linearly between the elevations that border the missing region. The aircraft's latitude, longitude, altitude, pitch and roll angle are also linearly interpolated during these conditions if the gaps are small (seconds) and the change in measured values before and after the gap is also small. All data points in the recorded waveforms above the aircraft altitude and below the ground return are removed. In the cases where there is no discernible ground return, points more than 200 m below the estimated surface elevation are removed.

Finally, the signal waveforms are vertically smoothed via a boxcar averaging window with the integration time equal to the 1-μs laser pulse width. Since the vertical resolution of the atmospheric backscatter measurements is primarily limited by the laser pulse width, the boxcar averaging has little effect on the vertical resolution of the backscatter signal, but it reduces higher frequency noise from the detector.

3.2 Level-1 Data Processing

The Level-1 data processing converts the waveforms from the Level-0 data processing to the attenuated atmospheric backscatter profile in terms of cross section per unit volume per steradian (m^{-1} sr^{-1}), also known as attenuated backscatter coefficient.

The optical signal power collected by the lidar can be written as,

$$y(t) = E_{tx} \int_0^{\tau_L} x(\tau)h(t - \tau)d\tau \tag{1}$$

where E_{tx} is the transmitted laser pulse energy in joules, $x(\tau)$ is the normalized laser pulse shape, i.e., $\int_0^{\tau_L} x(t)dt \equiv 1$, τ_L is the laser pulse width, and $h(t)$ is the impulse response of the atmosphere column, that is, the received signal in response to an infinitely short laser pulse. Here we assume the lidar detector has a much faster time response than the laser pulse width and the detector's output pulse shape is the same as that of the received optical signal.

The impulse response of the atmospheric column is related to the volume atmospheric backscatter coefficient profile by the lidar equation (Measures, 1984; Reagan et al., 1989)

$$h(t) = \frac{c}{2}\beta[R(t)]\,T_a^2[R(t)]\frac{\eta_r A_{tel}}{R^2(t)}C_1. \tag{2}$$

Here c is the speed of light, $\beta[R]$ is the backscatter coefficient at range R in units of $m^{-1}\,sr^{-1}$, $T_a(R)$ is the one-way atmospheric transmission from the lidar to range R, η_r is the lidar receiver optical transmission, A_{tel} is the collection area of the receiver telescope, $R(t) = c\,t/2$ is the lidar range at time t after the laser pulse emission, and C_1 is the lidar detector responsivity in V/W.

Substituting Eq. (2) into (1), the lidar signal from the detector can be written as,

$$s(t) = C_1 \frac{c}{2}\,\eta_r A_{tel} E_{tx} \int_0^{\tau_L} x(\tau)\frac{\beta[R(t-\tau)]T_a^2[R(t-\tau)]}{R^2(t-\tau)}\,d\tau. \tag{3}$$

If we assume that the atmospheric backscatter coefficient and the atmospheric transmission are approximately constant over the laser pulse interval (1 μs, or 150 m in range), the lidar signal from the detector can be written as,

$$s(t) \approx C_1 \frac{c}{2}\,\eta_r A_{tel} E_{tx} \frac{\beta[R(t)]T_a^2[R(t)]}{R^2(t)} = C_2 \frac{1}{R^2(t)}\langle\beta[R(t)]T_a^2[R(t)]\rangle \tag{4}$$

where $C_2 = C_1\eta_r A_{tel} E_{tx}\,c/2$ is a constant that dependents only on the parameters of the lidar. Therefore, the attenuated atmospheric backscatter coefficient as a function of range can be calculated from the lidar measurement by

$$\langle\beta[R(t)]T_a^2[R(t)]\rangle = [R^2(t)s(t)]\frac{1}{C_2}. \tag{5}$$

The lidar signal term $s(t)$ in the above equation can be expressed in terms of lidar range by substituting $t = 2R/c$. The attenuated backscatter coefficient can also be expressed as a function of altitude by subtracting the lidar range from the flight altitude of the aircraft.

For the CO_2 Sounder lidar configuration used in the 2017 airborne campaign, the instrument constants are

$$C_1 = \left(\eta_{det}G_{APD}\frac{1}{E_{ph}}e_c\right)Z_{TIA}G_{amp}\eta_{cable}, \quad \text{(V/W)} \tag{6}$$

$$C_2 = C_1\eta_r A_{tel}E_t\frac{c}{2} = C_1\,\eta_r\,\frac{\pi\phi_{tel}^2}{4}(1-L_{obs})E_t\frac{c}{2}. \quad \text{(V m}^3\text{)}. \tag{7}$$

Here η_{det} is the detector quantum efficiency, G_{APD} is the APD gain, $E_{ph} = h\,c/\lambda_{laser}$ is the laser photon energy, h is Planck's constant, λ_{laser} is the laser wavelength, e_c is the electron charge, Z_{TIA} is the gain of the preamplifier (a transimpedance amplifier) in V/W, G_{amp} is the post amplifier voltage gain, η_{cable} is the electrical cable transmission, ϕ_{tel} is the receiver telescope diameter and L_{obs} is the fractional loss of the telescope's collection area caused by its center obscuration.

Table 2 lists the values of these parameters for the 2017 CO_2 Sounder lidar. The resultant detector responsivity is $C_1 = 6.39\times10^8$ V/W for nominal APD gain, and the instrument constant is $C_2 = 5.13\times10^{10}$ V m^3.

Table 2. CO_2 Sounder Lidar parameters as used in the 2017 Airborne Campaign

Instrument Parameters	Values
Laser pulse energy, E_t	25 μJ
Laser pulse width, τ_L	1.0 μs
Laser pulse rate	10 kHz
Laser wavelengths, λ_{laser}	1572.2 to 1572.5 nm
Telescope diameter, ϕ_{tel}	0.20 m

Telescope center obscuration, L_{obs}	16%
Receiver optical transmission, η_r	81.3%
Receiver field of view, θ_{FOV}	500 µrad
Receiver optical filter width, $\Delta\lambda_f$	1.4 nm
Receiver integration time	1 s
Detector quantum efficiency, η_{det}	69.3%, including the fill factor
Detector avalanche gain, G_{APD}	190 (at 10 V APD bias)
Detector excess noise factor, F_{ex}	1.05
Detector noise equivalent power, NEP	1.7 fW/Hz$^{1/2}$
Receiver noise bandwidth for atmospheric backscatter calculations, B_n	0.5 MHz
Transimpedance amplifier gain, Z_{TIA}	320 kV/A
Post amplifier voltage gain, G_{amp}	12.6
Cable transmission, η_{cable}	90%
Overall receiver responsivity $(\eta_{det}G_{APD}/hc/\lambda_{laser})$ $Z_{TIA}G_{amp}(1-\eta_{cable})$	6.39e8 V/W at 10 V APD bias, decreases by a factor of two for every 1 V reduction in APD bias

The received signal level from the surface varies with lidar range. To prevent detector saturation, the gain of the lidar detector was adjusted manually during the flight to keep the received signal within the detector dynamic range. The detector gain was changed in steps of 2 by adjusting the bias voltage of the HgCdTe APD detector. In data processing, the detector's gain value was determined from the pulse amplitude of the window returns. After filtering out anomalous window returns, which generally resulted from the aircraft flying through clouds, the pulse energies and the centroid of the window returns are calculated and binned for each 1-s data file for each flight. The window return pulse energies are divided into groups separated by a factor of two. The average value from the bins with the APD biased at 10 V is defined as the nominal. Measurements made with window return pulse amplitude equal to ½ the nominal are assigned to have a gain correction factor of 2 and their backscatter waveforms are multiplied by 2. For those with window return pulse amplitudes equal to ¼ their backscatter waveforms are multiplied by a factor of 4, etc. When the window return pulse energy is an outlier or too weak to determine, the last known valid detector gain value is used.

For each second, the attenuated atmosphere backscatter coefficient is obtained by multiplying the signal waveforms from Level-1 data processing by the square of the lidar range and then dividing it by the lidar scaling factor C_2 according to Eq. (5). Figure 5 shows the atmosphere backscatter profile for the flight on 8 August 2017 along the flight path shown in Figure 6. Figures 7 and 8 show expanded views of the same data. During this flight aircraft flew over mountains, ocean, desert, etc. (Fig. 6) and performed seven spiral-down maneuvers to compare the lidar XCO$_2$ measurements with those from the in situ gas analyzer on board. The surface type and reflectance affected the solar background noise in the atmospheric backscatter measurement. For example, when the aircraft flew over ocean surface in between Spiral-Down #3 and #4, the scattered sunlight was much lower than that for land surfaces and the noise in the received data became noticeably lower, as shown in Fig. 5. Figure 9 shows a line plot of the vertical atmospheric backscatter profile measured at time 23:34:00 (UTC).

The raw lidar signal is digitized at 10-ns (1.5 m) intervals by the ADC to enable meter-level calculations of the lidar range to the surface. However, the vertical resolution of the computed atmospheric backscatter profiles is much wider due to the 1-μs (150 m) laser pulse width. To reduce the data volume, the backscatter profile data are re-sampled with 100 ns (15 m) bin width. This still oversamples the backscatter profile but preserves certain temporal features in the data, such as the ground returns and cloud boundaries. The horizontal resolution along the aircraft ground track is the distance travelled by the aircraft in 1-s, or about 200 m. These results from most flights showed the values of attenuated backscatter coefficients within the boundary layer are comparable to those reported by Spinhirne et al. (1997) which were measured at a nearby (1540 nm) laser wavelength. Similar atmospheric backscatter profiles have been measured with a ground-based 2-μm lidar that used the same type of HgCdTe APD detector (Refaat et al, 2020).

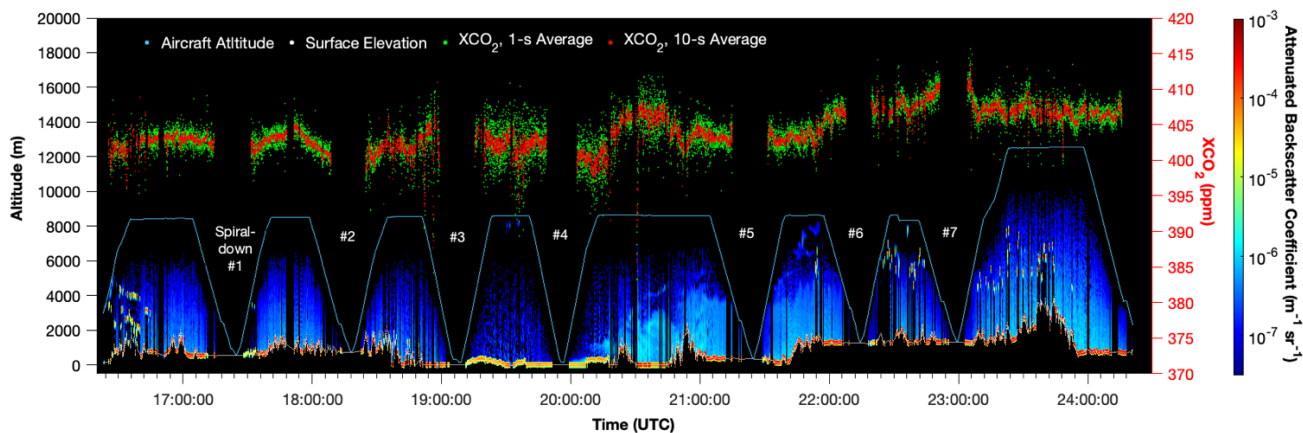



Figure 5. Time history of the attenuated atmospheric backscatter profiles measured by the CO_2 Sounder lidar during the flight on 8 August 2017 along with the with the retrieved values of XCO_2.

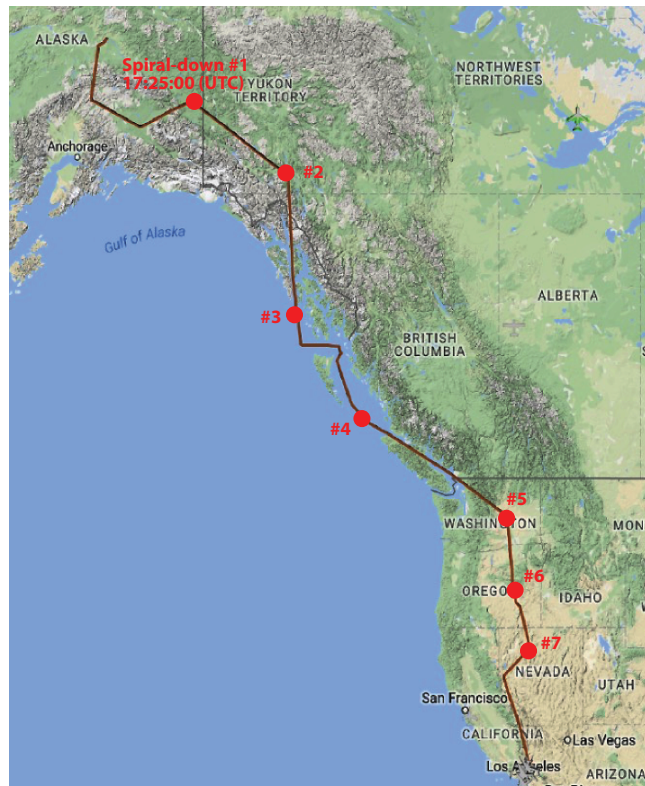



Figure 6. The flight path of the airborne CO_2 Sounder lidar on 8 August 2017 (plotted via Google Maps). Each dot indicates a spiral-down maneuver.

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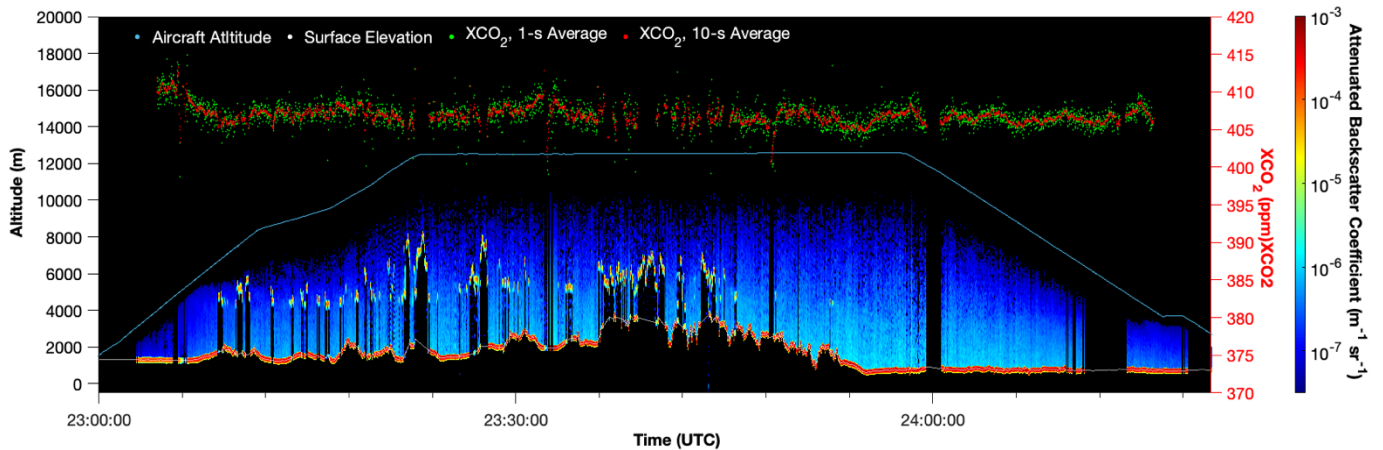
Figure 7. Same as Figure 5 but expanded for the last segment from time 23:00:00 (UTC) to the end of the flight.

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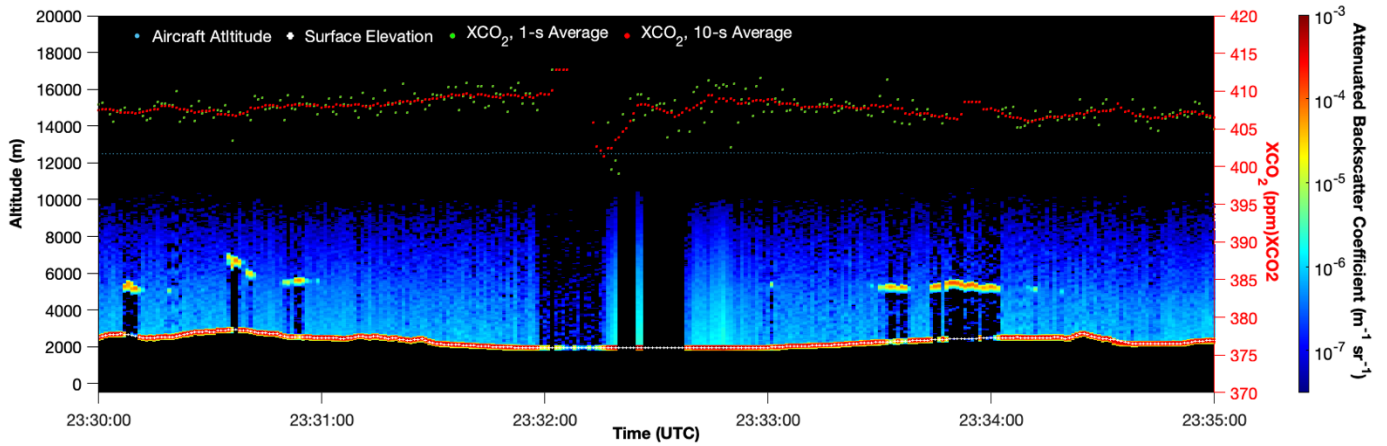
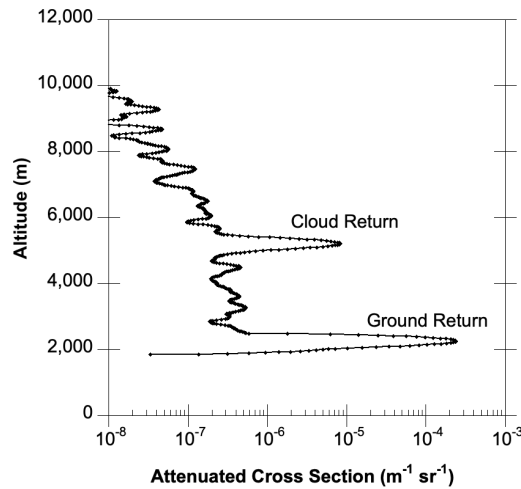


Figure 8. Same as Figure 7 but expanded for the 5 minutes between 23:30:00 and 23:35:00 (UTC).

320   **Figure 9.** Plot of the attenuated backscatter profile and ground return measured on 8 August 2017 at 23:34:00 (UTC) with 15-m vertical sampling interval. The results are from the same data shown in Figure 4 after Level 1 processing.

These backscatter profiles also contain the broadened lidar returns from the ground expressed in
325 terms of backscatter coefficient. With the 1.0-µs wide laser pulse and the 1.0-µs wide box-car smoothing, the smoothed return signal is usually triangular with a width of 10 range bins at the half maximum points and 20 range bins at the base. Any changes in surface elevation during the receiver integration time can cause some additional broadening of the return pulse shape to 15 to 20 range bins (1.5 to 2 µs), as shown in Figure 9. These ground return signals at the end of the backscatter profiles can
330 be used to calculate the attenuated surface reflectance by summing the backscatter values over all the range bins containing the ground return and then multiplying the results by π and the bin width (15 m).

Figure 10 shows the attenuated surface reflectance as the aircraft approached the Edward Air Force Base on 8 August 2017. Over a mountainous area for the time period from 23:00:00 to 23:40:00 (UTC), the attenuated surface reflectance varied between about 4% when there were clouds and 18% when cloud free. The average attenuated surface reflectance from 0:01:00 to 0:11:00 (UTC) over desert near the end of the flight was 22%. If we estimate the average desert reflectance to be 45% (Kuze et al., 2011), the one-way atmospheric transmission would be 70%, which seems consistent with the sky condition shown on the aerial photograph (Figure 11). When the surface reflectance can be estimated by some other means, the lidar profiles may be used to calculate the extinction to backscatter ratio of thin clouds and aerosols in the path.

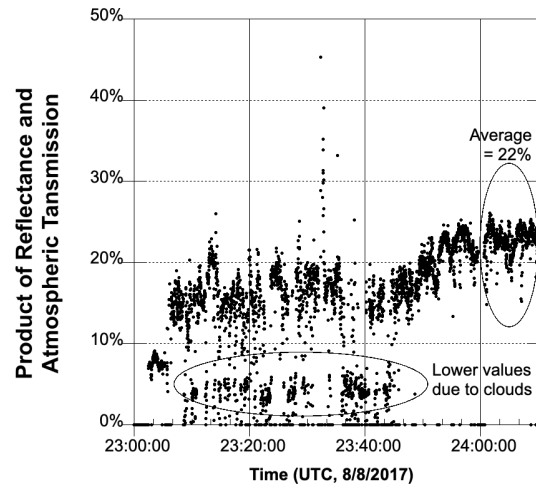



Figure 10. Product of the surface reflectance and two-way atmospheric transmission calculated by integrating the ground return of the backscatter profile as the aircraft flew toward the Edwards Air Force Base during 8 August 2017.

Figure 11. Photograph taken from the aircraft to the Edwards Air Force Base on 8 August 2017 24:14:00 (UTC) showing the sky condition for the last segment of data shown in Figure 10.

4. Signal to noise ratio calculations

To derive the SNR of the attenuated backscatter profiles we first express the average detected signal from each laser pulse in a way similar to Eq. (4), but in terms of the rate of detected signal photons as a function of the lidar range, as

$$\langle \dot{n}_{sig}(R) \rangle \approx \frac{c}{2} \frac{E_{tx}}{E_{ph}} A_{tel} \, \eta_{det} \, \eta_r \, \frac{\langle T_a^2(R) \rangle \langle \beta(R) \rangle}{R^2} \tag{9}$$

where R is the lidar range.

For each laser pulse, the standard deviation of the rate of the detected signal photons can be calculated as (Gagliardi and Karp, 1995),

$$\sigma_{noise}(R) = \sqrt{\left\{ F_{ex} \left[\langle \dot{n}_{sig}(R) \rangle + \langle \dot{n}_b \rangle \right] + \left(NEP \frac{\lambda_{laser}}{hc} \right)^2 \right\} B_n} \tag{10}$$

where F_{ex} is the detector excess noise factor from the randomness of the APD gain, $\langle \dot{n}_b \rangle$ is the average detected background photon rate, NEP is the detector noise equivalent power due to the detector dark noise and the preamplifier noise and B_n is the noise bandwidth of the lidar receiver used for the atmospheric backscatter calculations. The noise bandwidth is equal to $1/2t_{box}$ with t_{box} the width of the boxcar averaging window in the signal processing.

For the CO_2 Sounder lidar, the lidar range R in Eqs. (9) and (10) consists of a discrete set of lidar ranges for a series of range bins. Although the attenuated atmospheric backscatter profiles are given in the 15-m range bin size, the actual profiles are averaged over 150-m vertical layers due to the effect of the laser pulse width and the boxcar smoothing in the data processing. Therefore, the noise standard deviation given in Eq. (10) is for the 150-m vertical integration interval. This noise may be further reduced by averaging data from adjacent 150-m vertical atmospheric layers.

When the lidar views a sunlit diffuse scattering surface through a clear sky the average detected background photon rate is given by

$$\langle \dot{n}_b \rangle = I_s \, \Delta\lambda_f \pi \left(\frac{\theta_{FOV}}{2} \right)^2 \frac{\rho}{\pi} \cdot A_{tel} \, \eta_r \eta_{det} \tag{11}$$

where I_s is the solar irradiance at the 1572 nm laser wavelength, $\Delta\lambda_f$ is the receiver optical filter bandwidth, θ_{FOV} is the receiver field of view diameter, and ρ is the diffuse reflectance of the surface. These and other instrument parameter values are listed in Table 2.

For a single laser pulse, the SNR of the attenuated backscatter coefficient measurement is the same as the SNR of the rate of detected signal photons, which is the ratio of Eq. (9) to (10). Therefore, the SNR of a 150-m vertical bin at range R when averaged over the number of laser pulse measurements within the 1-s receiver integration time is given by

$$SNR_{av}(R) = \sqrt{N_{ave}} \, \frac{\langle \dot{n}_{sig}(R) \rangle}{\sigma_{noise}(R)} \tag{12}$$

with N_{ave} the total number of laser pulses measured within the receiver integration time. For the CO_2 Sounder lidar with a 1-s receiver integration time, $N_{ave} = 32 \times 9 \times 7 = 2016$.

The SNR is proportional to the attenuated backscatter coefficient and inversely proportional to the square of the lidar range. Hence for measurements from an airborne lidar, the SNR is a strong function of the range to the atmospheric layers being measured. Figure 12 shows the calculated SNR of the attenuated backscatter coefficient vs. lidar range for 10^{-6}, 10^{-5}, and 10^{-4} m^{-1} sr^{-1} backscatter coefficients at local noon and night. For example, for a 12-km aircraft altitude above ground surface and a 1-s receiver integration time, the averaged SNR is about 65 for clouds with a backscatter coefficient of
10^{-5} m^{-1} sr^{-1} that are 7 km below the aircraft (5 km above surface) during daytime at 45° sun angle. Therefore, these attenuated backscatter profiles may be used to identify mid-altitude clouds. The signals from aerosols at lower altitudes are much weaker. For example, at 10 km lidar range in the daytime, the averaged SNR is about 2.0 for aerosols of backscatter coefficient of 10^{-6} m^{-1} sr^{-1}. As a comparison, we calculated the average attenuated backscatter coefficient profile and the corresponding SNR calculated from 100 consecutive backscattering profiles starting from 8 August 2017, 23:55:00 (UTC) (a small cloud-free segment shown in Fig. 7). The mean and the ratio of the mean to standard deviation are plotted in Fig. 13. The ratio of the mean to standard deviation gives an estimate of the SNR. It is about 2.4 for attenuated backscattering coefficients of 10^{-6} m^{-1}sr^{-1} at 4.8 km altitude (7.8 km lidar range). The SNR would be 1.5 at 10 km lidar range. The SNR estimated from the measurement data is expected to be lower since it also includes the effects of small atmosphere variation during the measurement period. The SNR can be improved by averaging over longer time (more measurements along the flight path) and larger vertical layer thickness.

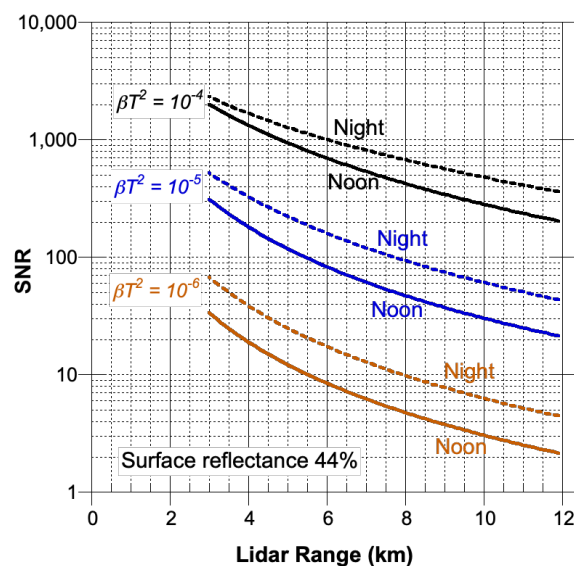



Figure 12. Estimated SNR of the attenuated backscatter coefficients obtained from the CO_2 Sounder lidar in the 2017 airborne campaign for 1-s integration times and 150-m vertical resolution. The SNR of the backscatter profiles may be further improved by averaging over a longer time intervals and by using thicker atmospheric layers.

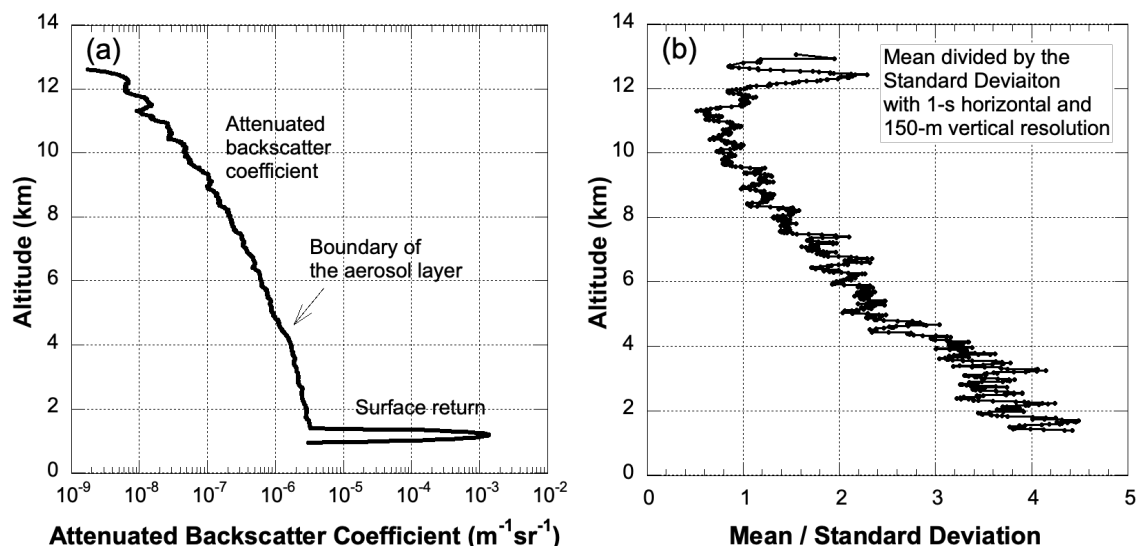



Figure 13. Average attenuated backscatter coefficients and the ratio of the mean to standard deviation calculated from 100 consecutive backscattering profiles starting on 8 August 2017 23:55:00 (UTC).

5. Conclusions

420   In addition to the measuring the column average CO_2 mixing ratio (XCO_2) the NASA GSFC CO_2 Sounder lidar measures the attenuated atmospheric backscatter profiles in the laser beam's path. We have recently processed the atmospheric backscatter profiles from the 2017 ASCENDS/ABoVE (Active Sensing of CO_2 Emission over Nights, Days, & Seasons /Arctic Boreal Vulnerability Experiment) airborne campaign and produced a new dataset of range-resolved attenuated atmospheric
425 backscatter coefficients at 1572 nm measured for each of the eight flights. The analysis shows the signal to noise ratios of the CO_2 Sounder lidar backscatter profiles are sufficient to identify clouds and to estimate the height of aerosol layers. This new data set also provides additional information to help interpret the retrieved XCO_2. These same types of measurements may be obtained in the future from a space-based lidar that uses the same measurement technique and has a similar value for (laser
430 power*telescope area/range^2).

Appendix

The following is the flowchart of the data processing software, which was developed using Matlab. The software consists of a main program "backscatterProcessing.m" and a set of functions indicated in blue text. The software code is available at the same website where the atmospheric backscatter profiles are
435 archived.

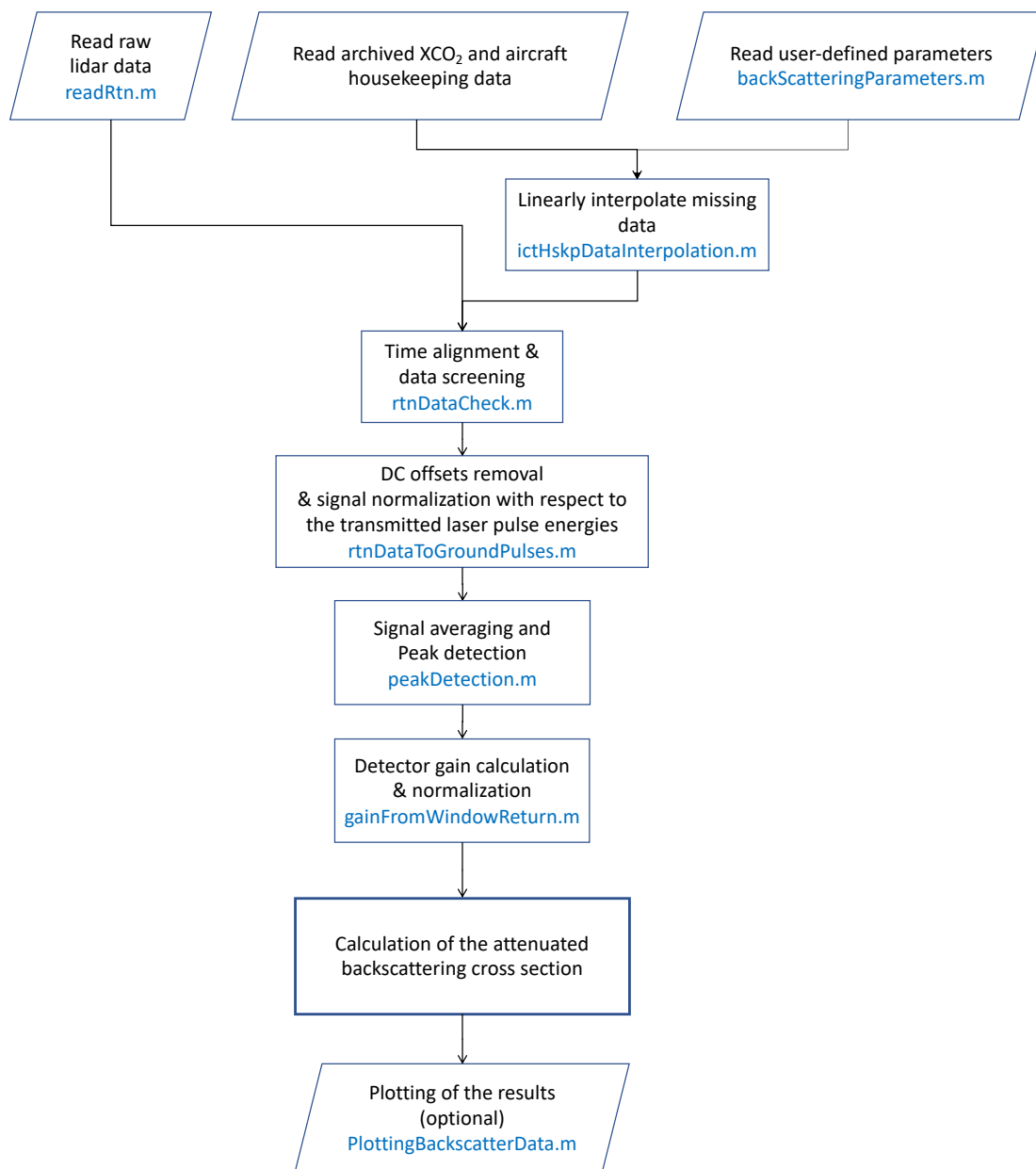



Figure A1. Flowchart of the software used to process the data described in this paper.

440 *Code availability.* The software codes used for the data processing of the atmospheric backscatter profiles described in this paper are posted on the same data depository website as the data.

Data availability. The atmospheric backscatter profiles by the CO_2 Sounder from the 2017 airborne lidar measurements are available from the NASA Oak Ridge National Laboratory (ORNL) Distributed Active
445 Archive Center (DAAC) for Biogeochemical Dynamics, https://doi.org/10.3334/ORNLDAAC/2051. NASA Earth Science Data are published under an arrangement equivalent to a CC0 licence. All data are openly shared, without restriction, in accordance with the EOSDIS Data Use Policy posted at https://earthdata.nasa.gov/earth-observation-data/data-use-policy?_ga=2.94803679.255626558.1646145252-1577112238.1614541217.
450

Author contributions. XS led the data processing and the writing of the manuscript; PTK developed the algorithm and the software to process the data; JBA led the CO_2 Sounder lidar team, the instrument development, the 2017 airborne campaign and its data analysis; SRK led the data archiving to the NASA Airborne Science Data website; and JM processed the 2017 airborne data for XCO_2 retrieval; all co-
455 authors participated the writing of the manuscript.

Competing interests. The authors declare that they have no conflict of interest.

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References

Abshire, J. B., Ramanathan, A., Riris, H., Allan, G. R., Sun, X., Hasselbrack, W. E., Mao, J., Wu, S.,
470 Chen, J., Numata, K., Kawa, S. R., Ying, M., Yang, M., and DiGangi, J.: Airborne measurements of CO_2 column concentrations made with a pulsed IPDA lidar using multiple-wavelength-locked laser and HgCdTe APD detector, Atmos. Meas. Tech., 11, 2001-2025, doi.org/10.5194/amt-11-2001-2018, 2018.
Allan, G., Abshire, J., Riris, H., Mao, J., Hasselbrack, W. E., Numata, K., Chen, J., Kawa, R.,
475 Rodriguez, Stephen, M.: Lidar measurements of CO_2 column concentrations in the arctic region of North America from the ASCENDS 2017 airborne campaign, Proc. SPIE, 10779, 1077906, doi.org/10.1117/12.2325908, 2018.
Gagliardi, R. M. and Karp, S.: Optical Communications, 2nd ed, John Wiley and Sons, 1995.
Kawa, S. R., Abshire, J. B., Baker, D. F., Browell, E. V., Crisp, D., Crowell, S. M. R., Hyon, J. J., Jacob,
480 J. C., Jucks, K. W., Lin, B., et al.: Active Sensing of CO2 Emissions over Nights, Days, and Seasons (ASCENDS): Final Report of the ASCENDS, Ad Hoc Science Definition Team, Document ID:

20190000855, NASA/TP–2018-219034, GSFC-E-DAA-TN64573, 2018, https://www-air.larc.nasa.gov/missions/ascends/docs/NASA_TP_2018-219034_ASCENDS_ID1.pdf.

Kuze, A., O'Brien, D. M., Taylor, T. E., Day, J. O., O'Dell, C. W., Kataoka, F., Yoshida, M., Mitomi, Y., Bruegge, C. J., Pollock, H., Basilio, R., Helmlinger, M., Matsunaga, T., Kawakami, S., Shiomi, K., Urabe, T., and Suto, H.: Vicarious Calibration of the GOSAT Sensors Using the Railroad Valley Desert Playa, IEEE T. Geosci. Remote, 49, 1781–1795, doi.org/10.1109/TGRS.2010.2089527, 2011.

Mao, J., Ramanathan, A., Abshire, J. B., Kawa, S. R., Riris, H., Allan, G. R., Rodriguez, M., Hasselbrack, W. E., Sun, X., Numata, K., Chen, J., Choi, Y., Ying, M., and Yang, M.: Measurement of atmospheric CO_2 column concentrations to cloud tops with a pulsed multi-wavelength airborne lidar, Atmos. Meas. Tech., 11(1), 127-140, doi.org/10.5194/amt-11-127-2018, 2018.

Mao, J., Abshire, J. B., Kawa, S. R., Haris, H., Sun, X., Nicely, J. M., Andel, N., and Kolbeck, P. T.: Measuring atmospheric CO_2 enhancements from the 2017 British Columbia wildfires using a lidar, Geophys. Res. Lett., 48, e2021GL093805, doi.org/10.1029/2021GL093805 2021.

Measures, R. M.: Laser Remote Sensing: Fundamental and Applications, Wiley-Interscience, New York, 1984.

Numata, K., Chen, J. R., Wu, S. T., Abshire, J. B., and Krainak, M. A.: Frequency stabilization of distributed-feedback laser diodes at 1572 nm for lidar measurements of atmospheric carbon dioxide, Appl. Opt. 50(7), 1047-1056, 2011.

Numata, K., Chen, J. R., and Wu, S. T.: Precision and fast wavelength tuning of a dynamically phase-locked widely-tunable laser, Opt. Express, 20, 14234-14243, doi.org/10.1364/OE.20.014234, 2012.

Refaat, T. F., Petros, Mulugeta, Remus, R., and Singh, U. N.: MCT APD detection system for atmospheric profiling applications using two-micron lidar, the 29th International Laser Radar Conference (ILRC'29), 2020, Article Number 01013, doi.org/10.1051/epjconf/202023701013.

Reagan, J. A., McCormick, M. P., and Spinhirne, J. D.: Lidar sensing of aerosols and clouds in the troposphere and stratosphere, Proc. IEEE 77(3), 433-448, 1989, doi 10.1109/5.24129.

Spinhirne, J. D., Chudamani, S., Cavanaugh, J. F., and Bufton, J. L.: Aerosol and cloud backscatter at 1.064, 1.54, and 0.53 μm by airborne hard-target-calibrated Nd:YAG/methane Raman lidar, Appl. Opt 36(15), 3475-3489, 1997.

Sun, X., Abshire, J. B., Ramanathan, A., Kawa, S. R. and Mao, J.: Retrieval algorithm for the column CO_2 mixing ratio from pulsed multi-wavelength lidar measurements, Atmos. Meas. Tech., 14, 3909-3922, doi.org/10.5194/amt-14-3909-2021, 2021.

Sun, X., Abshire, J. B., Beck, J. D., Mitra, P., Reiff, K., and Yang, G.: HgCdTe avalanche photodiode detectors for airborne and spaceborne lidar at infrared wavelengths, Opt. Express, 25(14), 16589-16602, doi.org/10.1364/OE.25.016589, 2017.

Sun, X., Kolbeck, P. T., Abshire, J. B., Kawa, S.R., and Mao, J., ABoVE/ASCENDS: Atmospheric Backscattering Coefficient Profiles, 2017. ORNL DAAC, Oak Ridge, Tennessee, USA. https://doi.org/10.3334/ORNLDAAC/2051 2022.