Long-term <u>ash dispersal dataset</u> of the Sakurajima Taisho <u>eruption</u> for <u>ashfall disaster countermeasure</u>

Haris Rahadianto^{1, 2}, Hirokazu Tatano², Masato Iguchi³, Hiroshi L. Tanaka⁴, Tetsuya Takemi², Sudip Roy⁵

¹Graduate School of Informatics, Kyoto University, Kyoto, 606-8501, Japan
 ²Disaster Prevention Research Institute, Kyoto University, Uji, 611-0011, Japan
 ³Sakurajima Volcano Research Center, Disaster Prevention Research Institute, Kyoto University, Sakurajima, 851-1419, Japan
 ⁴Center for Computational Sciences, Division of Global Environmental Science, University of Tsukuba, Ibaraki, 305-8577,
 Japan

⁵Department of Computer Science and Engineering, Indian Institute of Technology Roorkee, Roorkee, Uttarakhand, 247-667, India

Correspondence to: Haris Rahadianto (haris.rahadianto.88n@st.kyoto-u.ac.jp)

Abstract. A large volcanic eruption can generate large amounts of ash which can affect the socio-economic activities of

- 15 surrounding areas. Eruption scale and weather conditions primarily contribute to escalate ashfall hazards to wider areas. Accumulated ashfall has devastating impacts on both surrounding areas of the volcano and other regions, affecting airline transportation, socio-economics activities, and human health. Therefore, it is crucial to discover places with a high probability of exposure to ashfall deposition. In this study, we present an ashfall deposit and airborne ash concentration dataset from an ash dispersal simulation of a large-scale explosive volcanic eruption as a reference for ashfall disaster
- 20 countermeasures. We selected the Taisho (1914) eruption of the Sakurajima volcano, as our case study. This eruption regarded as the strongest eruption in Japan in the last century, and our study provides a baseline for the worst-case scenario. We employed one eruption scenario approach by replicating the actual eruption under various extended weather conditions to show how it would affect contemporary Japan. We generated an ash dispersal dataset by simulating the ash transport of the Taisho eruption scenario using a volcanic ash dispersal model and meteorological reanalysis data for 64 years (1958-
- 25 2021). We explain the dataset production process and provide the dataset in multiple formats for broader audiences. We <u>further clarified</u> the validity of the dataset with its limitations and uncertainties. By having an extensive dataset on ash dispersal in wider areas for a worst-case scenario, comprehensive countermeasure strategies can be derived to further reduce <u>ashfall risk</u>. The dataset is available at the DesignSafe-CI Data Depot: <u>https://www.designsafe-ci.org/data/browser/public/designsafe.storage.published/PRJ-2848v2</u> or through the DOI: <u>https://www.doi.org/10.17603/ds2-</u>
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70 1 Introduction

	A volcanic eruption is one of the events that emits several dangerous pollutants affecting human lives. Large eruptions eject
	enormous amounts of tephra and other eruptive materials into the air. Tephra hampers the residents near the volcano and Deleted: In the form of very fine ash, tephra hampers the ([1])
	citizens in the farther areas (Bonadonna et al., 2021) in the form of very fine ash. When a large eruption occurs, volcanic ash
	can travel far away from the volcano source, disrupting socio-economic activities in many different ways, for example,
75	damaging critical infrastructures and buildings, and causing health problems among the people (Wilson et al., 2012; Zuccaro
	et al., 2013). Settled volcanic ashes (i.e., ashfall) are directly affects human health and livelihoods destroying vegetation,
	crops, and pastures, causing physical damages to infrastructure, such as clogging drainage systems, contaminating water
	supplies, disrupting traffic, and damaging the vehicles on roads (Barsotti et al., 2010; Ayris and Delmelle, 2012; Damby et
	al., 2013; Zuccaro et al., 2013; Poulidis et al., 2018). The cumulative weight of ashfall on the roof <u>can</u> collapse buildings and
80	cause short-circuit electricity inside them (Zuccaro et al., 2013; Hampton et al., 2015). In addition, ash plumes from the
	volcanoes are a safety hazard that can severely damage commercial airlines (Folch et al., 2012; Peng and Peterson, 2012;
	Tanaka and Iguchi, 2019). The 2010 eruption of Eyjafjallajökull has amplified the evidence. During this event, volcanic Deleted: s
	ashes flew to thousands of kilometres away, forcing almost all airports in 25 countries to close for a week-long and costing a
	loss of billions of dollars in revenue for airlines (Folch et al., 2012; Peng and Peterson, 2012). Apart from the eruption scale,
85	the weather conditions, mainly winds, primarily contribute to escalating ashfall catastrophes to a broader range. <u>Ashfall</u>
	hazards are uncertain based on (1) the eruption magnitude and intensity and (2) the wind condition at the time of eruption
	that can bring ashes to distant places. Based on the evidence that ashfall hazards have dynamically changing exposures
	toward socio-economic activities, we believe that estimating the risk of finding effective countermeasures is critical.
	Currently, densely populated and modernised cities require comprehensive volcanic risk reduction strategies (Miyaji et al., Deleted: The presentCurrently, densely populated and modernised cities require comprehensive volcanic risk reduction strategies (Miyaji et al.,
90	2011). One of the necessary actions to lessen the impacts of ashfall hazards is assessing the risk to infrastructures, human
	lives, and economic impacts to develop a better response plan. Accordingly, the hazards and risk assessment for ashfall
	mostly relies on quantifying accumulated ashfall on the ground, then extending the analysis by combining social data, based ////
	on the interest (population data, building data, and many more). Many researchers have used this typical method to
	determine the expected impacts probability for future eruption events (Jenkins et al., 2014, 2015, 2018; Wilson et al., 2012; /
95	Biass et al., 2017). A comprehensive ashfall risk assessment requires both long-term weather and simulations for statistical
	analysis purposes, such as selecting typical conditions and analysing circumstances under exceptional conditions (Hattori et
	al., 2013, 2016). An important component of conducting such a study is to obtain long-term ashfall deposit data
	encompassing vast impact areas during an extended period. This paper presents the generated dataset of the ash dispersal
	products over a vast region <u>over 64 years (1958-2021)</u> to provide vital input <u>for the development of a more comprehensive</u>
100	ashfall risk assessment and emergency management. Sakurajima volcano was selected as an area of interest because of its
	high explosivity and potential for large-scale explosive eruptions within the following 20-25 years (Yamasato et al., 2013;
	Hickey et al., 2016; Biass et al., 2017; Poulidis et al., 2018).

	As supporting evidence, three large eruptions of Sakurajima volcano have occurred within the last six centuries: the Bunmei	~
	eruption in 1471, the An'ei eruption in 1779, and the Taisho eruption in 1914. The latter was considered the largest and the	
•	strongest eruption in Japan during the last century (Todde et al., 2017; Poulidis et al., 2018). Furthermore, Hickey et al.	
	(2016) showed that similar phenomena of magma supply rates that led to the Taisho eruption are currently occurring,	
155	suggesting 130 years return period from the last eruption to the next one in the future. Consequently, the Sakurajima volcano	
	tended to have another large-scale explosive eruption in the following multiple decades. To address, such an urgent issue, the	
	Kagoshima City municipal government has been preparing a risk management and evacuation plan by providing an updated	
	Sakurajima volcano hazard map. The hazard map contains the necessary information about the impacts of historical	
	eruptions, volcanic alerts and warnings, with guidelines on evacuation directions and procedures (Kagoshima City, 2010).	\sim
160	Furthermore, the Osumi Office of the Rivers and National Highways produced maps for the potential ash deposition from a	
	large eruption case in municipalities surrounding the Sakurajima volcano (Kyushu Regional Development Bureau, 2017).	
1	Recently, some studies have assessed the impacts of both continuous Vulcanian activities and scenarios for an explosive	
	eruption of Sakurajima volcano, providing insights into how volcanic ashes could disperse and affect livelihoods (Biass et	
	al., 2017; Poulidis et al., 2018). However, those studies and precautions only discussed the proximal impacts in short to	
165	medium terms, with no further analysis available for the impacts to the distal locations. Furthermore, the previous large-scale	1
	explosive eruptions have brought calamities across almost all of Japan. Thus, it is crucial to assess and develop a	· · ·
	comprehensive risk analysis for a broader range, because when the ashfall arrives in major socio-economic centres of Japan,	
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The daily record of ashfall deposit for an extended period (1958-2021), covering the southern Kyushu region, focusing mainly on Kagoshima prefecture as the closest urban area from the volcano.¶ Time-series record of airborne ash concentration, recorded at every one and a half hours after each simulation for an extended period (1958-2021).¶ (...[4])

240 This paper is structured as follows: Section 1, introduces the general background and motivation for this study, followed by Section 2, which presents the process and impacts of the ashfall dispersal from the Taisho eruption in 1914, derived from previous studies and historical reports of the eruption. Section 3 describes the methodology for producing the dataset, and Section 4 describes the format and the usage examples of the dataset. Section 5 thoroughly addresses the validation method with the limitations of the dataset, Finally, Section 6 concludes the paper. Unless otherwise specified, all time dimensions in this study use the Japan Standard Time (JST, +09UTC).

2 Ashfall dispersal during the Taisho eruption

The Taisho eruption was extensively studied by researchers all over the world, mainly due to:

- Detailed historical reports in English, were compiled and available shortly after the eruption (Omori, 1914; Koto, 1916).
- The concurrent unrest activities of <u>the Sakurajima volcano</u> (Iguchi, 2016; Iguchi et al., 2020; Poulidis et al., 2017, 2018).
 - 3. The high possibility of a large-scale explosive eruption in the <u>next decades that probably would</u> resembles the past event (Hickey et al., 2016).
- This section focuses on how volcanic ashes travelled during the eruption process, especially to distal locations spread across
 Japan. For reference regarding to the complete evidence and chronology of the Taisho eruption, please refer to the chronicles mentioned above and the studies conducted by Kobayashi (1982); Yasui et al. (2006, 2007); and Todde et al. (2017). The explanation here refers to, unless specified otherwise, both historical chronicles and the latest comprehensive study by Todde et al. (2017). The Taisho eruption is a large-scale explosive eruption with a Volcanic Explosivity Index (Newhall and Self, 1982) (VEI) of 4. The Taisho eruption ejected an enormous volume of eruptive materials from the two main active vents during its explosive phases. The explosive eruption started at 12 January 1914 10:00:00 and lasted 48 h. The tephra mostly ejected from the western vent (Nabeyama) mixed with those from the eastern vent (Yokoyama), simultaneously producing plumes estimated to be between 10 km and 18 km high. The westerlies at the upper altitude and the surface winds influence the ashfall dispersal pattern. Both the prevailing winds and plume height were significant drivers of the vast ash deposition process over a very wide area. Most tephra dispersed eastward, leaving Kagoshima City with only minor ash deposits during 265 the first several days of the eruption.

In contrast, the location close to the eastern vent had four meters deep ashfall deposit. Three days after the eruption started, the ashfall deposits were reported in Kyushu, Shikoku, western Japan, and Sendai in northern Japan. Village offices, tobacco plantations, and local meteorological observatories kept a record of the exact time of ash arrival and sighting, showing that volcanic ashes reached the major cities of Fukuoka at 08:00:00 and Osaka at midnight on 13 January 1914, Tokyo on the

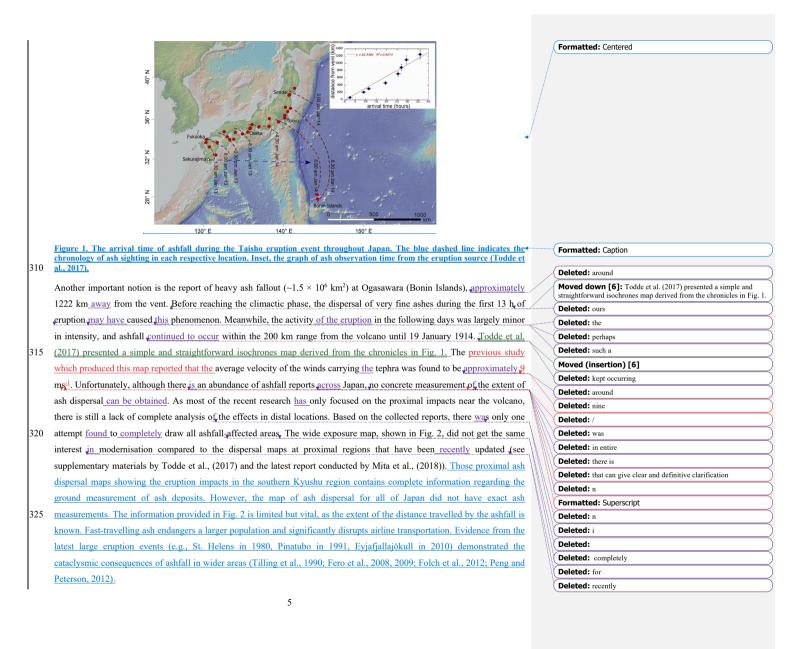
270 morning of 14 January 1914, and finally Sendai in the afternoon of the same day.

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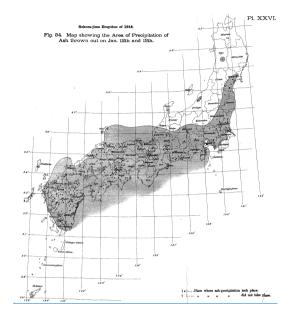


Figure 2. The ash dispersal map of the Taisho eruption on all of Japan (Omori, 1914),

Economic losses amounted to billions because obstructed air transportation was one of the more significant repercussions.
These losses were not realised by the authorities during the Taisho eruption event that occurred more than a century ago. Moreover, the Japanese airspace would suffer worse implications if a similar explosive eruption occurred shortly. At least 40% of the air traffic in the Japanese airspace will be affected if Sakurajima volcano has a large-scale explosive eruption in the present day. The number passengers, affected was almost four times higher than the number of people affected by devastating Typhoon Jebi back in 2018 (Takebayashi et al., 2021). The currently built environment and infrastructures 360 vulnerable to ashfall hazards require better comprehension of the upcoming volcanic risk for locations proximal as well as distal to the volcano. It is urgent to learn how large the catastrophe would affect contemporary Japan if the same event occurred in recent times.

Deleted: Those proximal ash dispersal maps showing the eruption impacts in the southern Kyushu region contains complete information about regarding the ground measurement of ash deposits. However, the map of ash dispersal for the entireall of Japan did not have an exact ash measurements at all. Though limited, the information provided in Fig. 2 is limited but vital, as the extent of the distance travelled by the ashfall travelled is Known. The Ffast-travelling ash endangers a larger population and greatly significantly disrupts airline transportation. Pieces of eEvidence from the latest large eruption events (e.g., St. Helens in 1980, Pinatubo in 1991, Eyjafjallajökull in 2010) demonstrated the cataclysmic consequences of ashfall ion wider areas (Tilling et al., 1990; Fero et al., 2008, 2009; Folch et al., 2012; Peng and Peterson, 2012)...

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having an extended period dataset on ash dispersal in wider areas for a worst-case scenario, hopefully, comprehensive countermeasure strategies can be derived to reduce the ashfall risk further.¶

3 Data generation methods

3.1 Simulation on selected eruption scenario

We used a fixed volcanological scenario over an extended contemporary period recorded in the meteorological reanalysis dataset to capture the extended daily variability of ash dispersal patterns from the Taisho eruption. We ran the PUFF model
for cases of 23,376 days from 1 January 1958 to 31 December 2021 The PUFF model is an ash tracking model developed during the Redoubt volcano eruption in 1989, which is famous for its high resolution and high accuracy results (Tanaka, 1994; Searcy et al., 1998; Scollo et al., 2011; Folch, 2012). This model considers ash as a collection of a finite number of virtual particles, computes the motion of the particles over time, and works better within the first 48 hours of the eruption

- (Searcy et al., 1998; Tanaka, 1994). The PUFF model is a Lagrangian-based model that has several advantages compared to
 other approaches (Eulerian), such as providing trajectory information, physical realism, describing, non-diffusive near-field
- to sources, numerical stability, lack of numerical diffusion, conservation properties, and resolution of sub-grid scale variability (Lin et al., 2012). Among all the nine Volcanic Ash Advisory Centers (VAACs) worldwide, eight VAACs, including Tokyo VAAC, use and operate models with the Lagrangian approach (Folch, 2012; Lin et al., 2012). A recent example of the use of the PUFF model for the simulation of explosive eruptions in Sakurajima volcano shows a satisfactory
- 405 outcome. This model <u>was</u> applied to cases of explosive eruptions in 2017 and 2018 to assess <u>its</u> utility in forecasting realtime ash fallout. Both simulations agreed well with the measurement, <u>which were</u> recorded by <u>the</u> instruments installed at several points nearby the volcano (Tanaka and Iguchi, 2019). <u>Here</u>, we construct the model using <u>many</u> random variables $r_{l}(t)$, where $i = 1 \sim M$ and M is the total number of particles. $r_{l}(t) = (x, y, z)$ represents the position vector for <u>the</u> *i*th particle at time *t* with its origin at the volcano vent. Using <u>the</u> discrete-time increment $\Delta t = 300$ s, the governing 410 equation can be written as:

$r_i(0) = S_i,$	$i = 1 \sim M$, for $t = 0$,
$r_i(t + \Delta t) = r_i(t) + V\Delta t + Z\Delta t + G\Delta t,$	$i = 1 \sim M$, for $t > 0$

where S_t is the initial location of all the particles at the vent, V = (u, v) is a vector for the wind velocity moving the particles and $Z = (c_h, c_h, c_v)$ is a vector for the diffusion velocity containing the diffusion speeds generated by Gaussian random numbers. $G = (0, 0, -w_t)$ is the gravitational fallout velocity <u>obtained</u> by approximating the extended Stokes <u>law for various</u> particle sizes (Tanaka, 1994; Tanaka and Yamamoto, 2002). The movement of particles steers the diffusion Z Z-direction, and the size of the particles affects the gravitational fallout *G* (Searcy et al., 1998; Tanaka, 1994). The diffusion speed *c* (c_h or c_v) was obtained by the random walk process related to the diffusion coefficient *K* as $c = \sqrt{2K/\Delta t}$ (Tanaka et al., 2016; Tanaka and Iguchi, 2019). Tanaka and Yamamoto (2002) conducted several diffusion tests with various values of *K*, and compared the results with satellite images of actual dispersals from several volcanic eruptions in the past. Based on this 420 research, we assigned a suitable horizontal diffusion coefficient $K_h = 150 \text{ m}^2\text{s}^{-1}$ and vertical diffusion coefficient $K_v = 1.5 \text{ m}^2\text{s}^{-1}$ as mentioned in Table 1.

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460 Table 1. Input values for the simulations used in this study.

Eruption Source Parameters	Value	
Mean Estimated Eruption Mass (× 10 ¹⁰ kg)	1.1	
<u>Mass Eruption Rate, ε (kg/h)</u>	5.8×10 ⁶ (Min) - 2.7×10 ¹⁰ (Max)	
Horizontal Diffusion Coefficient, K _h (m/s)	150	
Vertical Diffusion Coefficient, K _v (m/s)	1.5	
Log-Scale Mean Grain Size (mm)	1.0	

These diffusion coefficients are consistent with the in situ observations documented by Eliasson et al. (2014) and adjusted for the Sakurajima volcano. The values used in this study follow those reported by Tanaka et al. (2016, 2019) but are smaller than those reported by Fero et al. (2008, 2009) and Kratzmann et al. (2010). To further investigate the ash dispersal process to distal locations in entire Japan, we extended the simulation period of the PUFF model to 96 h after the eruption began.

- Following Eq. (1), by modelling the source location of the volcano vent as S(x, y, z), we can adjust the number of particles released at each time step to <u>obtain</u> the optimal statistical information from the model. For each time step, we assigned a Gaussian random number for each particle (M_0) and scaled it according to the emission rate, making the value of the total number of particles (*M*) changes linearly with the emission rate (Tanaka et al., 2016). We assigned the initial number of particles to 5000 due to constraints in the available computational power. The size of volcanic ash varies from fine ash to
- 470 boulders, as large particles tend to settle out <u>quickly</u> and then become smaller with time, Because of this process, each ash particle will have a <u>different</u> grain size (Bonadonna et al., 1998), and we assume that each particle has an initial grain size of a logarithmic Gaussian distribution with a standard deviation of <u>2.0</u> centred at <u>-3.0</u>. Thus, the average particle size was approximately 1.0 mm on a log scale, with 68% of the particles are within the range of 10 <u>µm_10</u> cm, as shown in Table 1. Given the initial vertical velocity of the emission with specified damping (e-folding) time τ_0 , the particles are distributed 475 randomly in a vertical manner from the vent z_1 to the plume's peak z_2 continually. Then, using time integration for the

vertical velocity in the momentum equation, we obtained the final form as $z_0 = z_1$ and $z_1 = z_2$ during the time step Δt (Tanaka et al., 2016),

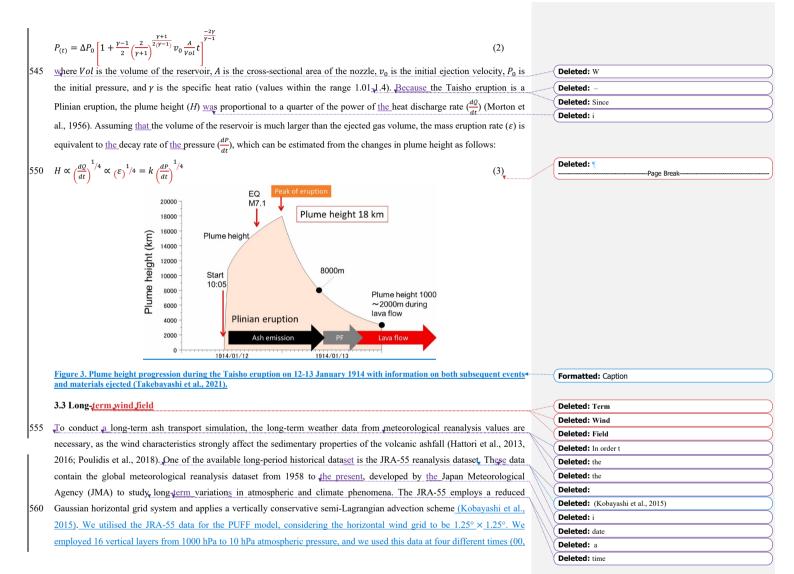
3.2 Estimation on mass eruption rate based on plume heights transition

In this study, we used the estimated emission rate from the calculation done by Iguchi (2014), based on a previous study (Kobayashi, 1982). As shown in Fig. 3, we consider the maximum estimated plume height at 17,890 m around midnight on 13 January 1914 which is close to the range of plume heights in the chronicles explained in Section 2. We obtain the value of mass eruption rate (ε) and the mean eruption mass (see Table 1) with the help of temporal change in plume height. The obtained values agreed well with a recent observation (Todde et al., 2017). The pressure of the reservoir that ejects gas through a nozzle (P_(t)) at time t decays exponentially with time (Nishimura, 1998) following Eq. (2):

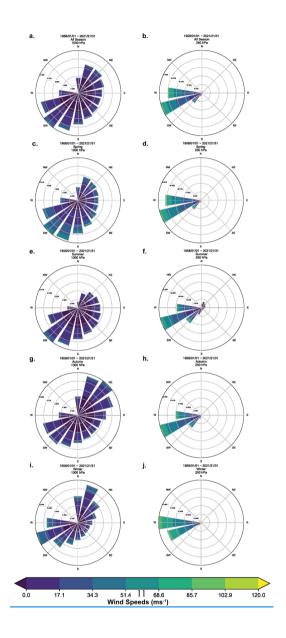
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<u>06, 12, and 18 UTC</u>) daily. The 3D wind field consisted of zonal wind (*U*), meridional wind (*V*), and geo-potential height (*gph*), which were interpolated to the position of each ash particle using the cubic spline method in space and time.



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- Figure 4. The ten wind roses illustrate the wind conditions inferred from the JRA-55 Reanalysis dataset (Kobayashi et al., 2015)
 for both surface winds (first column) and upper altitude winds at approximately 10 km asl (second column). The first row ((a) and (b)) portrays all wind conditions for the entire 1958-2020, while the rest shows the influence of seasonality. (c) and (d) are the wind conditions in spring, (e) and (f) are the wind conditions in summer, (g) and (h) are the wind conditions in autumn, and (i) and (j) are the wind conditions in winter. Roses correspond to the probability of the provenance and sequential colours indicating the speed gradients.
- 590 In Fig. 4, each wind rose corresponding to the probability of the provenance identified by sectors, and sequential colours mark the speed gradients. As discussed in Sect 2, the ashfall dispersal from the Taisho eruption was controlled by both surface winds (1000 hPa), and westerlies (250 hPa), We only portrayed the wind conditions at those pressure levels, with the first column indicating the upper altitude winds and the second column indicating the latter. The second cow onwards, reflects seasonal variations: spring (March to May), summer (June to August), autumn (September to November), and winter
- 595 (December to February), whereas the first <u>row</u> displaying <u>all the</u> dates selected for the simulation (1958-2021). As noted by previous research concerning the climatological conditions in the Sakurajima area, wind conditions change heavily <u>depending</u> on the season. Thus, the ash deposition process largely depends on the time of the eruption (Biass et al., 2017; Poulidis and Takemi, 2017; Poulidis et al., 2018)

3.4 Recording ashfall Jocation and thickness

- 600 This study considers the ashfall deposits on the ground as particles with non-positive altitudes, with their location marked as a longitude-latitude pair. For each simulation, as the computation <u>progressed</u>, the particles with a negative value in the altitude dimension were subset to other files before being compiled when the simulation <u>was completed</u>. During this process, we captured the rest of the flowing particles as airborne ash concentrations, copied them to different files, and saved them using time marking every 90 minutes. We separated the measurement of the ashfall deposit and airborne ash concentrations
- 605 into two different files using this mechanism. Further<u>more</u>, we <u>measured</u> the ashfall thickness (ashfall depth) by dividing the surface area <u>by</u> the ashfall density according to its particle size. First, we allocated the ashfall particles to grids according to / their location. Then, we set all ashfall particles in <u>those grids according to their</u> average values. We considered all the / adjacent grids and the pre-assigned mass values multiplied by the number of particles to obtain the total mass of the ashfall deposit. It is possible to obtain the mass of each particle based on the particle size. However, because we deliberately
- 610 designate the total initial number of particles as a constant number and smaller than the actual number in the real case, the total mass of all particles is less than the actual <u>mass</u>, and the difference may increase as the scale of the eruption increases. Therefore, considering the particle size under the total mass conservation, the virtual mass of each particle is necessary (Shimbori et al., 2009). When the eruption reached its peak, each ash particle contained approximately 450 tons of virtual mass. This method was performed to alleviate the uncertainty that appeared when assigning a small number of particles in
- 615 the simulation (see Scollo et al., 2011). After assigning the virtual mass to each particle in one grid, the thickness of the ashfall χ_{ij} at grid *i*, *j* obtained from the following equation:

 $\chi_{ij} = \sum_{n} \frac{m(D_n)}{\Delta x \Delta y}$

(4)

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where *m* is the total virtual mass for the total *n* particles within the grid *i*, *j* of area $\Delta x \Delta y$ corresponding to the particle size D_n (Shimbori et al., 2009). All computations from running the simulation, processing wind fields, and computing the total ashfall accumulation on the ground were performed with a small workstation with 32GB RAM, dual 16-cores Intel Xeon E5-2620V4 2.10 GHz processors, and 4GB NVIDIA Quadro P1000 GPU.

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4. Dataset <u>overview</u> and <u>usage examples</u>		Deleted: Overview
4.1 Data description and availability		Deleted: Availability
.1 Data description and availability		Formatted: Heading 2
The dataset contains three data:		Deleted: Each ash dispersal simulation produces
1. Ashfall deposit measurements from 128.5° E to 148.6° E, and 30.0° N to 45.9° N in 0.1°×0.1° grid (~10 km ²) for an		
extended period (1958-2021).		

- Ashfall deposit measurements from 129° E to 132° E, and 30.5° N to 32.5° N in 0.01°×0.01° grid (~1 km²) for an extended period (1958-2021).
- The time series collection of the total airborne ash concentrations in the air was recorded for every 90 minutes for an extended period (1958-2021).
- 720 To ensure maximum availability for the diverse users of the dataset, we prepared the dataset in two formats. We stored the dataset as a space-separated ASCII value table in comma-separated value (CSV) format (see Table 2) and a multi-dimensional array structure in the Network Common Data Format (NetCDF). NetCDF is useful for supporting access to diverse types of scientific data, and its files are self-describing, network-transparent, directly accessible, and extensible (Unidata, 2021). We developed the NetCDF files using the *Xarray* library in Python using the NETCDF4 package (Hoyer

715

725 and Hamman, 2017). Owing to the high dimensionality of the airborne ash concentrations, we only provided these data in the NetCDF format. Meanwhile, the ashfall deposit data are available in CSV and NetCDF formats.

Table 2: The tabular view of the sample deposit data for both regions. Each column is separated by a comma, corresponding to the <u>CSV format</u>.

<u>Deposit/CSV/Japan/<yyyy>/<yyyymmdd.csv></yyyymmdd.csv></yyyy></u>			Deposit/CSV/Kag	oshima/ <yyyy>/<y< th=""><th><u>YYYMMDD.csv></u></th></y<></yyyy>	<u>YYYMMDD.csv></u>
latitude	longitude	<u>deposit</u>	<u>latitude</u>	<u>longitude</u>	<u>deposit</u>
<u>30.00</u>	128.50	<u>0.00</u>	<u>30.00</u>	129.00	<u>0.00</u>
<u>30.00</u>	128.60	<u>0.10</u>	<u>30.00</u>	<u>129.01</u>	<u>0.10</u>
<u></u>	<u></u>	<u></u>	<u></u>	<u></u>	<u></u>
<u>30.10</u>	128.50	<u>0.20</u>	<u>30.01</u>	<u>129.00</u>	<u>0.20</u>
<u>30.10</u>	128.60	<u>0.30</u>	<u>30.01</u>	<u>129.01</u>	<u>0.30</u>
<u>%.2f</u>	<u>%.2f</u>	<u>%.2f</u>	<u>%.2f</u>	<u>%.2f</u>	<u>%.2f</u>

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We set the name of each dataset file as the simulation date in ISO-8601 format (*YYYYMDD*<*<.csv*, .nc>>), stored inside a directory named by year (*YYYY*) within each respective region (Japan and Kagoshima). In addition, each dataset contains an ordered set of location markings (longitude, latitude) and ashfall depth (in cm), which are stored in two decimal floatingpoint formats. The set of location markings is the coordinates in the NetCDF data format of the ashfall deposit data. In <u>contrast</u> to the deposit data, the collection of airborne ash concentrations provides the total number of particles at a specific location and time. The tracking period of the ash particles started 90 minutes after the simulation and was completed after 96

- 740 h, As the three-dimensional shape of the affected area and the total number of particles change throughout the time dimension, we assign *date time* as the primary coordinate. The dataset directory consists of two main folders based on the data observations: *Airborne* and *Deposit*. The *Deposit* folder contains two child directories based on the data format and observation range, and then breaks down to every year starting from 1958 and ending in 2021. The *Airborne* dataset has a straightforward directory (year, *YYYY*) because it only consists of one data format, and there is no separate observational
- 745 range. The detailed structure of the dataset for both the ground deposits and airborne concentrations is provided in the Appendix A.

The dataset is available at the DesignSafe-CI Data Depot (https://www.designsafe-ci.org) hosted at the Texas Advanced Computing Center (TACC). This data depot is provided by the Natural Hazards Engineering Research Infrastructure (NHERI), which provides the natural hazards engineering community and researchers with comprehensive state-of-the-art

750 cyberinfrastructure (Rathje et al., 2017). The dataset can be accessed directly at <u>https://www.designsafe-ci.org/data/browser/public/designsafe.storage.published/PRJ-2848v2</u> or through the DOI: <u>https://www.doi.org/10.17603/ds2-ww5f-t920</u> (Rahadianto and Tatano, 2020). Users can access the dataset directly without any prior registration and <u>may</u> choose any data to <u>be</u> downloaded. The data can be downloaded as a single file or as a collection <u>of multiple files. Please</u> select the <u>version 2</u> of the dataset for the latest updates.

755 4.2 Usage examples of the dataset

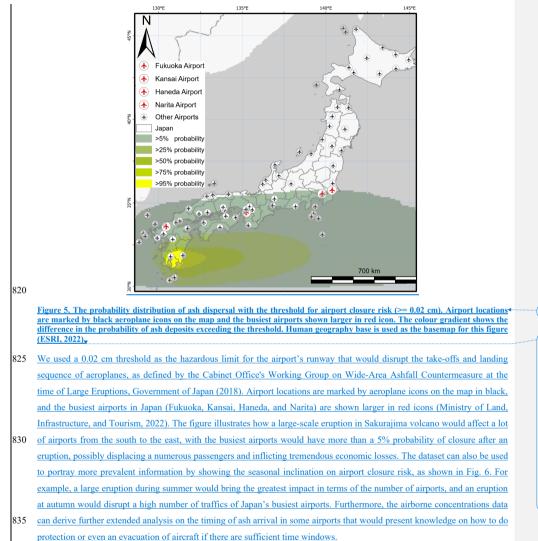
The dataset comprises ash deposits on the ground and airborne concentrations of a typical large-scale eruption at Sakurajima volcano (Sect. 2) for an extended period. Given the size of the dataset, the users can use the data to prepare a response for an ashfall disaster. Here, we demonstrate how to use these data to create conditional ashfall hazard maps. The information in such maps can serve as a reference for designing an early warning system, planning ash clean-up operations and closure

- 760 management on air and land transportation, and even preparing for long-term long-distance evacuation strategies. The maps describe the probability distribution of exceeding a selected hazard threshold based on the statistical distribution of the wind profiles (Bonadonna, 2006). Several researchers used such maps to illustrate hazard assessments of a specific volcanic eruption scenario, identify the possible impacts of the assumed maximum expected events, and quantify the probability of first-order impacts on built environments (Bonadonna, 2006; Biass et al., 2014, 2016). Figure 5 reveals the conditional
- ashfall hazard maps accounting for the airport closure risk of all of Japan for all the wind profiles used in the simulation.

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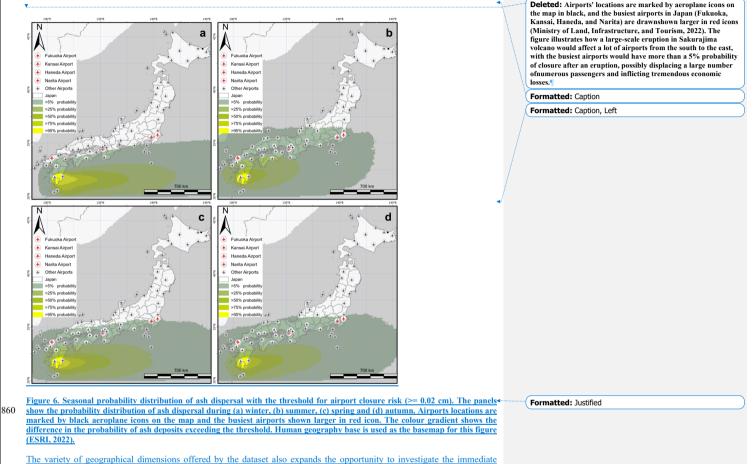
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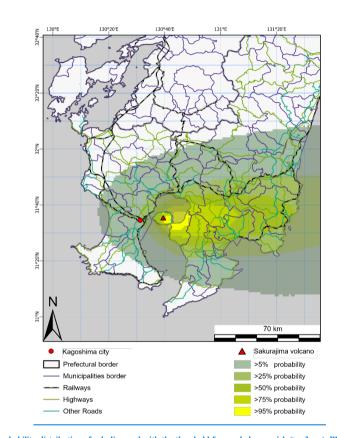
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consequences of ground ashfall deposits in the proximal urban areas surrounding Sakurajima volcano. Figure 7 displays the conditional ashfall hazard maps accounting for the airport closure risk of all of Japan for all the wind profiles used in the simulation. We use a 3 cm threshold as the hazardous limit for the closures of roads and railways, as it would make them impassable by cars and trains, as defined by the Cabinet Office's Working Group on Wide-Area Ashfall Countermeasure at the time of Large Eruptions, Government of Japan (2018).

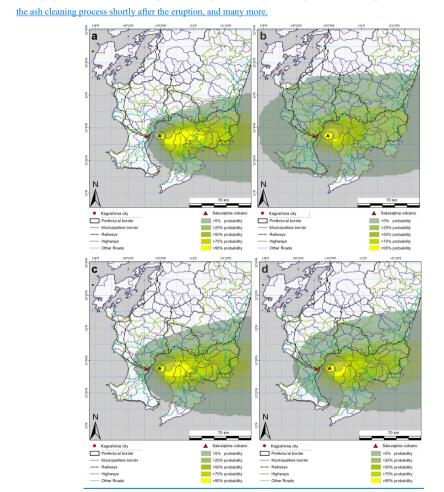


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Figure 7. The probability distribution of ash dispersal with the threshold for road closure risk (>= 3 cm). Black and white stripes mark railways on the map, and highways and other roads are in coloured lines. The colour gradient shows the difference in the probability of ash deposits exceeding the threshold. Human geography base is used as the basemap for this figure (ESRI, 2022).

885 The figure illustrates how a large-scale eruption in Sakurajima volcano can severely affect the eastern parts of Kagoshima City, located downwind of the volcano with more than 25% probability, regardless of eruption time. The data also show how the seasonal patterns determine the extent of the ash deposits in the western parts of South Kyushu, with a summer eruption escalating the damage to almost the entire Kagoshima Prefecture, as depicted in Fig. 8. The intensity of the ashfall hazards to the transportation networks can provide necessary insights for preparing responses and countermeasures.





890 These preparations include scheduling roads and highway blockages, making evacuation strategies for safer areas, managing

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Knowing tThe intensity of the ashfall hazards to the transportation networks can giveprovide necessary insights tofor preparing responses and countermeasures. These preparations, which include scheduling roads and highway blockages, making evacuation strategies tofor safer areas, managing the ash cleaning process shortly after the eruption, and many more....

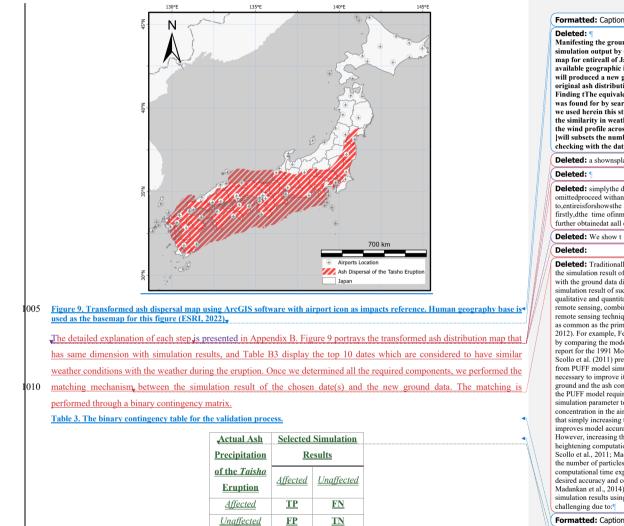
Figure 8. Seasonal probability distribution of ash dispersal with the threshold for road closure risk (>=3 cm). The panels show the probability distribution of ash dispersal during (a) winter, (b) summer, (c) spring and (d) autumn. Black and white stripes mark railways on the map, and highways and other roads are in coloured lines. The colour gradient shows the difference in the probability of ash deposits exceeding the threshold. Human geography base is used as the basemap for this figure (ESRI, 2021)_x

	5 Validation and Jimitations	S
	5.1 Validation_ <u>result</u>	
935	Section 3.1 briefly describes the strength and the well-known performance of the PUFF model for ash dispersal simulations.	
	Several researchers validated their simulation results through various methods, such as direct comparison to post-event	
	reports, matching the ash clouds trajectory with satellite images, and adjusting the model's input to find the best-fit values	\mathbb{N}
	(Fero et al., 2008, 2009; Folch, 2012; Scollo et al., 2011; Tanaka and Iguchi, 2019; Tanaka et al., 2020; Webley and Mastin,	
	2009). In addition, several past studies have validated ash dispersal simulations using the PUFF model for multiple eruption	
940	cases in Sakurajima volcano, albeit in smaller-scale eruptions (Tanaka and Iguchi, 2019; Tanaka et al., 2020).	
	For historical large-scale eruptions of Sakurajima volcano, which had larger plume height, eruption volume, and wider range	
	of ash deposition, the validation has never been done. Jdeally, if the precise wind data at the time of the selected eruption are	
	available, we can replicate the eruption; thus, simulation results with similar wind patterns of the eruption would resemble	N
	the historical reports. The dataset presented here offered an opportunity to perform the validation of ash dispersal simulation	$\backslash \rangle$
945	of a large-scale eruption of Sakurajima volcano over a wide space (all of Japan). We employed the conventional method by	
	directly comparing the selected simulation results with the available reports of the Taisho eruption. However, in this study,	
	making a direct comparison with the ground truth was challenging owing to the following:	·
	1. Lack of complete ground reports. Section 2 outlined the available reports gathered in cities and tobacco plantations	
	across Japan only mentioned subjective ash deposit observations without a quantitative measurement, e.g., "frost-	
950	like ash", "house roof become white", and "slight deposit",	
	2. Complete wind data at the time of eruption were absent. Only surface observation are available at limited	
	observation points, and it is not possible to completely replicate the ash transport process,	
	Despite these limitations, we attempted to utilise an innovative approach to present a validation result using the dataset	
	presented here with all available information. The ash dispersal map presented in Fig. 2 provided the locations of the ash	
955	deposition to all of Japan is the. Moreover, the detailed weather conditions that asserts control over the ash dispersal process	
	are available as well. Therefore, we utilised to these information to proceed with the validation process, as follows:	Annual States
	1. Manifesting the ground data in the same dimension as the simulation output by digitally redrawing the ash	
	distribution map for all of Japan, provided by Omori (1914), using available geographic information software	
	(GIS). This procedure produced a new ground observation data that resembled the original ash distribution of the	
960	Taisho eruption.	
	2. The equivalent wind pattern on the day of the eruption was found for similar weather on the wind data we used in •	
	this study, JRA-55 Reanalysis. We expected that the similarity in weather features would also bear a similarity in	
	the wind profile across multiple atmospheric levels. This process subsets the number of simulation results for	
	validation checking with the data produced in the previous step,	

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Manifesting the ground data in the same dimension withas the simulation output by digitally redrawing the ash distribution map for entireall of Japan, provided by Omori (1914), using available geographic information software (GIS). This procedure will produced a new ground observation data that resembled the original ash distribution of the Taisho eruption.

Finding tThe equivalent wind pattern on the day of the eruption was found for by searching the similar weather on the wind data we used herein this study, JRA-55 Reanalysis. We expected that the similarity in weather features would also bear a similarity in the wind profile across multiple atmospheric levels. This process lwill subsets the number of simulation results for validation checking with the data produced in the previous step.

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Deleted: Traditionally, the primary way to confirm the accuracy of the simulation result of volcanic ash transport models is to compare with the ground data directly. Another method to validate the simulation result of such models is through satellite imagery for both qualitative and quantitative measurement. Apart from the space-based remote sensing, combining multiple ground and aeroplane-based remote sensing techniques can validate the models, although it is not as common as the primary method (Webley and Mastin, 2009; Folch, 2012). For example, Fero et al. (2009) measured the model's accuracy by comparing the model output with satellite data and the deposit report for the 1991 Mount Pinatubo (Webley and Mastin, 2009). Scollo et al. (2011) presented a novel mechanism to validate results from PUFF model simulations by evaluating the number of particles necessary to improve its results for the tephra deposit value in the ground and the ash concentration in the air. The study claimed that the PUFF model requires up to 33 and 220 million particles set in the simulation parameter to match the field data for ash deposits and their concentration in the air, respectively. This finding further confirms that simply increasing the particle number set in the model input improves model accuracy (Scollo et al., 2011; Bursik et al., 2012). However, increasing the number of particles in the simulation means heightening computational requirements (Peterson and Dean, 2002; Scollo et al., 2011: Madankan et al., 2014). Furthermore, increasing the number of particles in the simulation means increasing the computational time exponentially, requiring a trade-off between desired accuracy and computational cost (Scollo et al., 2011; Madankan et al., 2014). In this study, we found that validating our simulation results using the aforementioned conventional ways is challenging due to: ... [23]) Formatted: Caption

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Table 3 briefly explains the contingency table used to evaluate the simulation results. The evaluation focused on measuring how the ash dispersal map produced from the simulation on selected dates would equally cover an equivalent region of interest to the ground data.

135 We limited the evaluation for all ash deposits on the land, and omitting all ash deposit on the sea. Then, by matching the binary contingency table, we quantified the hit rate score *Hit* of the simulation result as follows:

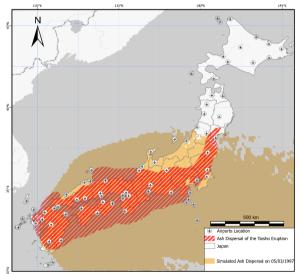
 $Hit = \frac{TP + TN}{TP + TN + FP + FN}$

(4)

<u>*TP*</u> denotes the true positive, which means that both the ground data and the simulation results have an ash deposit on the same grid points (affected). *TN* denotes the true negative, in which the ground data and the simulation results don't have an

140 ash deposit on the same grid points (unaffected). Meanwhile, *FN* denotes the false negative, in which particular grid points are marked as unaffected in the simulation results but appear to be affected in the ground data. In reverse, *FP* denotes the false positive, in which particular grid points are marked as affected in the simulation results but appear to be unaffected in the ground data. Therefore, the hit rate score *Hit* is the simulation result correctly dispersed ashes to the same grid points as the ground data, divided by all defined grid points from the original ash dispersal map. To make it brief, we chose only the top date of the STS scores list in Table B2 (5 March 1997) in the matching process. The simulation results of the chosen date

reached satisfactory performance with a high *Hit* score (0.832). This result means that the selected date with weather conditions similar to those of the Taisho eruption have a considerable identical ash dispersal pattern.



Deleted: Table 3. The binary contingency table for the validation	
process	

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Figure 10. The comparison of the ash distribution coverage between the original ash dispersal map and the simulation results onthe selected date (5 March 1997). The red-stripped area indicates the original areas affected by the Taisho eruption, and the light orange colour represents the ash dispersal from the simulation. Human geography base is used as the basemap for this figure (ESRI, 2022),

Figure 10 further exhibit the evidence of the similarities in the ash dispersal map for the selected simulation results. The process to produce such a result involve a complex procedure (detailed in Appendix B) because of the limited availability of the necessary components for the direct comparison. Because of this complication, such a process may not yield perfect results, both for the checking components (new ground data and selected wind patterns due to similar weather conditions)

and the matching ash distribution of the simulation.

5.2 Dataset limitations

Generally, the ash deposition results from the simulation of volcanic ash dispersal and transport models rely on how the parameters values of both eruption source parameters and the wind conditions during the eruption are set. These parameters are critical in determining the impacts of ashfall on proximal and distal locations (Bonadonna et al., 2012; Folch, 2012; Macedonio et al., 2016; Mastin et al., 2009; Webley and Mastin, 2009). This study used the deterministic value of plume height, erupted mass, and mass eruption rate as the eruption source parameters. We utilised the calculation by Iguchi (2014), incorporating the study by Morton et al. (1956) for the relationship between the observed plume height in the historical

- 170 reports and studies (Kobayashi, 1982; Koto, 1916; Omori, 1914; Todde et al., 2017) to produce <u>the</u> mass eruption rate. <u>There</u> was a positive correlation between the change in plume height over time and both the total eruption mass and mass eruption rate. However, <u>we advise caution</u> when using this method, as it introduces a significant bias in the mean height value. The eruption rate is roughly proportional to the fourth power of height. <u>Therefore</u>, error in assigning the precise height will <u>significantly</u> affect its conversion to <u>the</u> eruption rate (Folch et al., 2012; Mastin et al., 2009). We assign the parameters of
- 175 the last eruption, as a baseline of the largest eruption size or plume height as a conservative safety factor, although we recognise wide variability that could change the significance of the implied hazard. The values used in this study agreed well with the past observations, and the bias presented in the eruption source parameter should be small and would not significantly affect the ash dispersal mechanism. However, variations in these parameters can result in different eruption rates and total mass, resulting in different ashfall footprints. In addition, future vent location, eruption style and size, and
- 180 eruption duration significantly affect the impacts of the ashfall in both proximal and distal areas (Bonadonna et al., 2012; Mastin et al., 2009; Selva et al., 2018) Furthermore, depending on the model, the parameters which need to be set arbitrarily, such as the diffusion coefficient, settling particle velocity law (see Table 1), orographic effects, and the lack of topographical data, could cause discrepancies if different variations were used (Macedonio et al., 2016; Poulidis et al., 2017, 2018; Scollo et al., 2011). Finally, the distribution of the total grain size, the particle density and shape, and the chosen aggregation
- 1185 method also affect the accuracy of the model outputs (Bonadonna et al., 2012; Folch et al., 2010, 2012; Mastin et al., 2009; Selva et al., 2018). Here, we approach the simulation with the main purpose <u>of characterising the Taisho eruption</u>, which eliminates the variations in those variables <u>because</u> we used a fixed scenario as <u>the worst case</u>.

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Moved up [13]: Meanwhile, *FN* denotes the false negative, in which particular grid points are marked as unaffected in the simulation results but appear to be affected in the ground [24]

Deleted: Meanwhile, FN denotes the false negative, in which particular grid points are marked as unaffected in the simulation results but appear to be affected in the ground data. In reverse, FP denotes the false positive, in which particular grid points are marked as affected in the simulation results but appear to be unaffected in the ground data. Therefore, the hit rate score Hit is the simulation result correctly dispersed ashes to the same grid points as the ground data. Therefore, the hit rate score Hit is the original ash dispersal map. To make it brief, we chose only the top date of the STS scores list in Table B2 (5 March 1997) in the matching process. The simulation results of the chosen data reached satisfactory performance with a high Hit score (0.832).... This result means that the selected date with weather conditions similar to those of the Taisho eruption have a considerable identical ash dispersal naterms.

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Deleted: 5.2 Finding the Identical Weather Characteristics The validation method, as shown in Fig. 9, consists of tw(....[26] Deleted: the most critical parts ...n determining the how ashfall impacts of ashfall on proximal and distal locations (Bonador(....[27])

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Deleted: we used in this study agreed well with the past observations, and the bias presented in the eruption source participation [30].

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We acknowledged the limitations of this study, which used the fixed values for the eruption source parameters to precisely replicate the transition in the Taisho eruption, which strongly restricts the various possibilities of future events. Following the result mentioned earlier by Hickey et al. (2016), which illustrates how the current uplift situation is similar to that of the preceding eruption event, our results could provide valuable input to both authorities and stakeholders. The dataset presented here will provide a guideline for an insightful understanding of how such a large explosive eruption would affect contemporary Japan. Moreover, inaccuracies can occur due to the effects of the winds and atmospheric humidity, among other factors (Folch et al., 2012; Mastin et al., 2009). Meteorological conditions determine the ashfall dispersion trajectory

- and ashfall deposit location (Macedonio et al., 2016). Therefore, reanalysis datasets are favourable for use because they have better accuracy than the forecast ones (Folch et al., 2012). It should be noted that the variability in the different meteorological databases (e.g., ERA-Interim, JRA-55, NARR, etc.) is relatively insignificant (Selva et al., 2018). As intended to provide guidance for a worst case scenario, the dataset provided can remain helpful in extending the present study toward the ashfall risk analysis and its decision-making process. We argue that the dataset here is vital to support
- 505 future researchers and emergency managers in analysing and planning crisis responses <u>under different conditions</u>. It is / important to note that these assignments are rough estimates of one of the possibilities <u>for the most likely future eruption size</u> / and type at Sakurajima volcano. We understand that there might be inconsistencies in the dataset that likely introduces a bias in the ashfall hazard and risk analyses. Future users should consider all these sources of uncertainty.

6 Conclusion

- 1510 Ash dispersal products from large eruptions have devastating impacts on various sectors, both in areas near the volcano and in distal locations. Emergency managers and crisis response planners need to acknowledge the compounded effect of ashfall deposits on the infrastructures to manage losses appropriately. This study presents a dataset that is useful for more extensive / research and planning, focusing on the context of the entire country. This paper describes the complete generation process of a dataset based on the recent large eruption of Sakurajima volcano (the Taisho eruption) by replicating the eruption process
- 515 over an extended period from 1958 to 2021. Furthermore, we added a validation result of ash dispersal simulation of a largescale eruption of Sakurajima volcano over a wide space (all of Japan) using the selected simulation results from the dataset, The complications and constraints of the validation procedure have been acknowledged, Finally, we acknowledge, the limitations and uncertainties of the dataset, Although it contains some degree of inaccuracies, the dataset can still encourage further studies in the ashfall risk analysis and decision-making process as limited as possible. Such advancements will be result to support further studies in the ashfall risk analysis and decision-making process as limited as possible. Such advancements will be
- 520 crucial to support future researchers and emergency managers in devising disaster responses <u>under different conditions</u>.

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Appendix A

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The structure of the dataset in the NetCDF format are shown in Fig. A1 for ground deposits data and Fig. A2 for airborne

630 concentrations data.

Deposit/NetCDF/Japan/<YYYY>/<YYYYMMDD.nc> Deposit/NetCDF/Kagoshima/<YYYY>/<YYYMMDD.nc> ordinates Coordinates latitude (latitude) float32 30.0 30.1 30.2 ... 45.7 45.8 45.9 latitude (latitude) float32 30.0 30.01 30.02 ... 32.99 33.0 8 longitude eruption_date longitude (longitude eruption_date () (longitude) float32 128.5 128.6 128.7 ... 148.5 148.6 9 (longitude) float32 129.0 129.01 ... 131.99 132.0 2 0 datetime64[ns] ... datetime64[ns] ... 9 3 Data variables: • Data variables: (latitude, longitude) float32 ... deposit (latitude longitude) float32 deposit . Attributes: • Attributes: Name : Long-period Ashfall Deposition Dataset (1958-2021) from the Sakurajima Taisho E Name : Long-period Ashfall Deposition Dataset (1958-2021) from the Sakurajima Taisho E ruption. Daily Astrial Deposit for Kagoshima region from simulation on 1958/01/01 10:00 .57 (01/UTC). This dataset presents the ashfall deposit thickness in cm for Japan. The dataset ge neared from a numerical simulation of the 1914 Jaisho eruption in Sakurajima us ing PUFF model and JRA-S3 Reamilysis data. ruption. Daily Ashfall Deposit for Japan region from simulation on 1958/01/01 10:00 JST (0 Type : Type : 1UTC). This dataset presents the ashfall deposit thickness in cm for Japan. The dataset ge Description : Description : nerated from a numerical simulation of the 1914 Taisho eruption in Sakurajima us ing PUFF model and JRA-55 Reanalysis data. Edition : Edition -Authors : Publisher License : Haris Rahadianto Authors : Publisher : License : . Waris Rahadiante Haris Rahadianto Disaster Prevention Research Institue (DPRI), Kyoto University, Japan Open Data Commons Attribution License (ODC-By) Disaster Prevention Research Institue (DPRI), Kyoto University, Japan Open Data Commons Attribution License (ODC-By)

Figure A1. The structures of the sample deposit data for both regions in the NetCDF format.

Airborne/<YYYY>/<YYYYMMDD.nc>

Coordinates:				
date_time	(date_time) da	tetime64[ns]	1958-01-01T11:30:00 1958-01	
eruption_date	0 da	tetime64[ns]		. 2
Data variables:				
longitude	(date_time)	float32	888	
latitude	(date_time)	float32		
altitude	(date_time)	float32		
ash_particles	(date_time)	float32		
Attributes:				
Name :	Long-period As ruption.	hfall Deposit	ion Dataset (1958-2021) from the Sakurajima	Taisho E
Type :	Airborne Ashes y 90 minutes.	for 96 hours	from 1958/01/01 10:00 JST (01UTC), recorded	at ever
Description :	simulation date	. The dataset	tal number of airborne ash particles from the s generated from a numerical simulation of the a using PUFF model and JRA-55 Reanalysis dat	e 1914 T
Edition :	1 .		-	
Authors :	Haris Rahadiant			
Publisher :			h Institue (DPRI), Kyoto University, Japan	
License :	Open Data Con	nmons Attrib	ution License (ODC-By)	

Figure A2. The structures of the sample airborne data in the NetCDF format.

1635 Appendix B

The diagram explaining the flow of the validation procedure for this study is portrayed in Fig. B1

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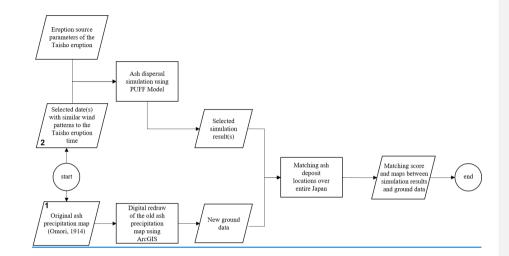


Figure B1. The flowchart of the validation procedure using the dataset produced in this paper. (1) the process of digitising the original ash precipitation map to become new ground data, and (2) the process of selecting a simulation result for matching with the ground observation data.

B.1 Transforming the original ash precipitation map

In the first step, as shown in Fig. B1, we transformed the ash dispersal map in Fig. 2 using ArcGIS software with a WGS 84 projection (ESRI, 2017). In addition, we added the current airport locations to signify the impacts of the very fine ashfall on entire Japan. We conducted the map transformation by placing binary markings to indicate the affected regions, i.e., 1 for red-stripped grids and 0 for the rest. All the binary values have a pair of longitude-latitude values to represent their positions on the map. This procedure is necessary, as we finally have a ground data in the same area dimension as the model output to directly compare it with the selected simulation result.

B.2 Finding the identical weather characteristics

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After we obtained the new ground data, the next steps matching a simulation result with it. However, we cannot simply check the data with any results from the simulation because of the unusual weather at the time of the eruption. Due to the absence of a complete wind profile at the time of the eruption, we investigated the weather conditions on that day. Understanding the weather components at the time of eruption is crucial for determining the behaviour of the ash fallout trajectory. This feature of the Taisho eruption is often omitted, although it explains why the ashfall from the eruption moves erratically to distal locations. Several prior researchers did not proceed with their analysis of this anomaly, which has proven to be influential concerning the impacts in distal municipalities (e.g., Biass et al., (2017), among many others). Formatted: Heading 2

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The Taisho eruption occurred in winter when mid-latitude cyclones were active. With the stationary planetary waves generated by the land-sea contrasts and the flow over the Himalayas, these phenomena reinforce the tropospheric jet stream 1660 and the baroclinic wave to be the strongest. In Japan, the reinforced jet streams cause the westerlies at 250 hPa pressure to exceed 70 ms⁻¹. The winds in the tropospheric jet stream blow from the west throughout the year; they are strongest in winter and weakest in summer (Wallace and Hobbs, 2006). The Meteorological Research Institute of the JMA (JMA-MRI) further clarifies how the westerly wind traverses above the Sakurajima volcano all year long from west to east. The westerlies reach their peak velocity in January and then continue to weaken until their lowest speed in summer, before rising again (Maeda et al., 2012). This cycle runs for the entire year and is one of the deciding factors for the ash dispersal pattern.

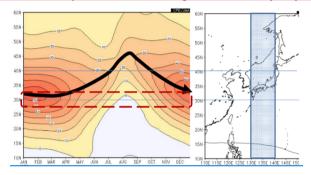


	Figure B2. Seasonal progression of the westerlies over Japan at 12 km from the surface (Maeda et al., 2012). The red dashed	(Formatted: Caption	
	square indicates the region covering Sakurajima volcano and Kagoshima city. The vertical axis shows the latitude degree of the			
1(70	wind position. The horizontal axis shows monthly changes of the westerlies velocity for the left image and the longitude of the			
1670	winds on the right image. The thick black arrow from left to right indicating the location of maximum velocity of westerlies based			
	on seasonal progression. The straight blue lines indicate all regions in Japan affected by the westerlies. The small number inside			
	the left image is the wind contour.			
	Figure B2 shows a time-latitude cross-section of the east-west wind at the 250 hPa (approximately 12 km from the surface),		Deleted:	
	averaged at 130°E to 140°E based on the 30-year survey (Maeda et al., 2012; Mita et al., 2018). The Sakurajima volcano and			
1675	its vicinity are situated around 31° N (red dashed square), including Kagoshima City. Therefore, the westerlies would bring			
	most of the ashfall from the eruption to the east. Instead, chronicles reported that ashes traversed to the north, west, and			
	southwest (Omori, 1914; Todde et al., 2017). This pattern contrasted with the preceding eruption (An'ei eruption -			
	November 1779) that followed the pattern of the westerlies (Tsukui, 2011). Research investigating the past ash dispersal		Deleted: ¶	
	patterns from Sakurajima volcano also reaffirms the differences due to seasonal features, agreeing that during winter, tephra			
1680	should mostly go eastward, with most of it falling in the Pacific (Biass et al., 2017; Poulidis et al., 2018; Mita et al., 2018).			
	When the eruption occurred, the vicinity of Kagoshima City and Sakurajima volcano experienced a maximum atmospheric_		Deleted:	
	pressure the first four days, providing a unique characteristic of the Taisho eruption.			
•				

From 9 January 1914 to the morning of 13 January 1914 atmospheric pressure in Japan consistently recorded an unusually high, resulting in stable sea waves and low-velocity winds, indicating calm and clear weather in most places (Omori, 1914). Unexpectedly, the weak winds helped many people living near the volcano, evacuated by boat (Kitagawa, 2015). The eruption seemed to play a significant role in increasing the temperature in Kagoshima, but apparently, it turned out to be the 1690 opposite (Omori, 1914). The leading cause of the high temperature at the time of eruption time was a sudden increase in southerly winds, as observed in Okinawa, 373 km from the vent (Omori, 1914). Furthermore, the chronicles explained the tri-daily general weather conditions during the eruption and found that similar weather phenomena also occurred on 17 January 1914 (Omori, 1914). This vital information helped us derive a method for validating our simulation results. Following such reports, we concluded that clear, sunny weather with weak winds played a major role in transporting ashfall 695 to entire Japan, and such conditions, although rare, can be repeated. From this observation, we assume that the wind pattern will also be similar on days with identical weather characteristics as irregular as the day of the Taisho eruption. Using this assumption, we further obtained wind data at various atmospheric pressure heights matching the wind data during the eruption. Thus, the simulation on the dates with similar weather is expected to produce a similar ash distribution to all of Japan, comparable with the transformed ash dispersal map in Fig. 9. 700 Owing to seasonal differences, finding dates with similar weather only uses winter dates, because the Taisho eruption occurred in winter as well. In detail, we searched and found the day(s) with weather conditions like the eruption day (12 January 1914). The search process only focused on winter days (December, January, February, and March) from 1 January 1958 to 31 December 2021. The important main features to query are clear sunny days, weak winds with high atmospheric pressure (anti-cyclonic), and higher temperatures, which are quite distinct and rarely occur in Japanese winter. These features 705 reiterate the peculiarity of the Taisho eruption and its diverging ash distribution. Weather comparisons were conducted using available weather information from the chronicles and historical weather information provided by the JMA on their website. However, the past meteorological data are limited because the Japanese authority only maintained surface weather condition (temperature, pressure, humidity, precipitation, and wind speed) in a major meteorological observatory located in the capital of selected prefectures. These past data consist of daily and hourly observations. Apart from surface observations, there are 710 no other available measurement recorded at different atmospheric pressure levels. Other available information is the surface weather chart on the day of the eruption (12 January 1914) and all simulation dates. Therefore, we decided to utilise a surface weather chart instead of surface observations, as it provides a clearer description of a weather phenomenon on a particular day. A weather chart, also known as a weather map, usually consists of symbols and features that describe a particular meteorological pattern in a specific space-time dimension. Weather charts were created by plotting the relevant 1715 measurements such as mean sea-level pressure, wind barbs, and cloud cover. This plotting can help find synoptic-scale features, such as weather fronts. A meteorologist usually performs an isobaric analysis of these maps, which involves the construction of lines of equal mean sea level pressure (Wallace and Hobbs, 2006).

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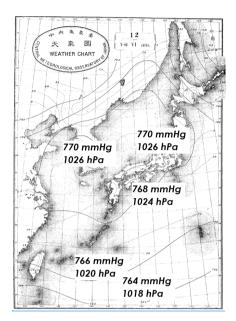


Figure B3. The weather chart on the Taisho eruption with additional markings of pressure measurement converted to the latest convention (mmHg -> hPa). Modified from Omori (1914).

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725 Figure B3 shows the weather chart at 06:00:00 on 12 January 1914, with the modern convention of pressure measurements. This weather chart displays how the high-pressure condition covers almost the entire Japan centring the central Honshu. These local maxima in the pressure field, denoted by a big "H" letter, can simply explain the common weather characteristics in that specific region. The wind becomes stronger when the pressure gradient between the high- and low-pressure systems increases. Land friction weakens the wind coming out from high-pressure systems. Hence, high-pressure systems typically 730 result in weaker winds with clear skies and relatively warm weather (Wallace and Hobbs, 2006). These implications

corroborate the weather reports during the eruption day. Accordingly, using the similarity on the surface weather chart is appropriate for identifying identical weather conditions.

B.2.1 Weather report commentary on the surface weather charts

Our initial attempt to search for identical weather by using weather chart similarity is by employed an image similarity method to compare features drawn in a pair of weather charts. However, this attempt turned out to be very difficult, as the weather chart on the eruption day was much different from the modern weather chart currently used in Japan.

τ		ometric pressure, the dime	ension and projection o	f the map, and th	e area included in	Deleted:	
the map were a		current daily weather chart		*		Deleted: 1	
40 the daily weath	her charts in Japan has c	hanged several times after	World War II. Starting	g in 1883, the JM	A produced daily		
surface weathe	er charts for the Asia-Paci	fic region, consisting of Ja	pan and Japanese-occu	pied areas pre-W	orld War II. Since		
August 1958,	the JMA has provided da	aily weather charts at vario	ous upper levels (500 l	nPa, 700 hPa, and	850 hPa) for the		
Asia-Pacific ar	nd 500 hPa for the North	ern Hemisphere region. Th	en, starting in March 1	996, the JMA add	ed a daily surface		
weather chart	for the Northern Hemisp	here and a 300 hPa level f	for the Asia-Pacific reg	gion. Finally, in F	ebruary 1999, the		
15 JMA updated a	all the weather charts and	added a specific surface w	veather chart for the reg	gions of Japan. Th	ese changes were		
also accompar	nied by changes in the w	eather chart format, and t	he current daily weath	er charts follows	the 1999 format.		
Thus, owing to	o the ever-changing form	ats of past weather charts	and the modern ones,	performing imag	e similarity is not		
feasible and is	a time-consuming task b	ecause we need to have mu	ultiple training data for	every format to d	ecide whether the		
two images ar	re similar. We moved to	another option which inc	ludes employing other	information avai	lable in the daily		
50 weather chart.	Another piece of information	ation available inside the w	eather chart is the gene	eral weather repor	t commentary (天		
<u> 気概況</u>). Expe	erts in meteorology develo	pped this commentary as a	guide to weather condi	tions during a part	icular day. Figure		
B4a portrays th	he weather report comme	ntary on 12 January 1914 lo	ocated at the bottom lef	t of the second pa	ge.		
a.	すり	東京地方南東,風屋雨探揉了り道:北東,風晴後,底方雪東北世道東北地方、南東,風景九州南部四國內海地方、南東,風景丁雨北環球、北東,風景九州南部四國內海地方、南野風雨後,晴九州北琉球、北東,風景九州南部四國內海地方、南野風雨後,晴九州北		HOUSE AND			

Figure B4. The general weather report commentary on the weather condition inside the weather chart. (a) The old weather chartput the commentary on the bottom of the second page on the day of the Taisho eruption (Database of Weather Charts for Hundred Years, 2022). (b) The modern weather chart mentions the commentary on the bottom of the monthly compilation of daily weather charts (Japan Meteorological Agency, 2022).

- All historical daily weather charts, from 1883 up to March 1938, have a commentary written on a weather chart explaining 760 the weather conditions. However, this feature was omitted from April 1938 until December 1995. The JMA resumed adding a general weather report commentary feature to the daily weather chart from January 1996 to date, under the weather chart images in the monthly compilation of historical daily weather charts (Database of Weather Charts for Hundred Years, 2022). Figure B4b shows the first page of the daily weather charts compilation for December 2020 from the JMA as a reference
- 765 (Daily Weather Chart, 2022). The commentary usually includes the important weather characteristics on that day, such as the existence of high- or low-pressure centres or fronts, the description of extreme weather phenomena such as typhoons, heavy rains, or heavy snow, and other special events on the day, such as disasters. This description allows us to find similar weather between the dates used in the simulation and the day of the eruption, even though we further decreased the number of dates used in the searching process (1996-2021, 3120 days). However, before proceeding with the searching process, we need to
- 770 perform several conversions because there is a minor distinction between the old commentary and the modern ones, which are:

1. The old commentary on the weather chart on the day of the eruption was based on old Japanese syntax.

2. The old commentary addressed both the weather in Japan and the weather in a Japanese-occupied areas.

- We solve these slight differences by simply converting the old commentary to contemporary Japanese syntax and removing all observation areas outside the modern Japan region. The first column in Table B1 denotes the original form written in the 775 Old Japanese syntax. The second column shows the transformation from the Old to the contemporary Japanese. The old commentary consists of three parts: explaining the location of high- or low-pressure centres, the weather on the main islands of Japan (Honshu, Kyushu, and Hokkaido), and lastly on other parts of the Asia-Pacific, coincidentally, at that time all regions mentioned are under Japanese occupation. This commentary agrees with the meteorological report provided in the 780 chronicles (Omori, 1914). Therefore, we omitted the last phrases and focused only on the weather on the main islands of Japan. Then, further converted it to a form similar to the commentary in the current daily weather charts. The new phrases in
- the last column give more brief comments that match with the contemporary style in the modern weather charts while still maintaining the important features (high-pressure all over Japan, clear/sunny weather in main islands, Kyushu is cloudy, and Hokkaido is either clear or cloudy with strong south-westerly winds). Then, we compared these phrases in contemporary
- 785 Japanese with all the commentaries found in the modern weather chart (from 1 January 1996 to 31 March 2021) provided by the JMA in winter (from January 1996 to March 2021), from the historical archive collected by the National Institute of Informatics (Database of Weather Charts for Hundred Years, 2022). These commentaries explain the weather conditions on a particular day in a random sequence, that is, not in a similar order to the processed commentary phrases (from 12 January 1914). Finally, we utilised the natural language processing (NLP) method to find semantic textual similarity between all 1790 commentaries with sentence embedding using SentenceBERT (Reimers and Gurevych, 2019).

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Table B1: The conversion of the general weather commentary in the weather chart on 12 January 1914.

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Original Form	Contemporary Form	Final Form	Formatted Table
高気圧東に移り、中心本邦中部に	高気圧は東に移動し、日本の中心にあ		
<u> 在りて、773mm を示し。揚子江</u>	<u>り、気圧計の値は 773mm を示すして</u>		
低気圧は目下漢口付近に存す。本	いる。長江にある低気圧は、ちょうど	<u>日本付近は移動性高気圧に覆わ</u> <u>れている。日本列島わ晴れ。一</u> <u>方、九州は曇り。北海道は南西</u> 風が吹き、晴れや、曇り。	
<u>邦にては、天気おおむね良好なれ</u>	武漢付近に存在する。日本では、天気		
<u>とも、九州にてはすでに曇天とな</u>	は晴れているが、九州地方では既に曇		
<u>り、北海道にては南西風強く晴曇</u>	りとなっており、北海道では南西の風		
<u>相半す。満州および朝鮮にてはお</u>	が強く晴れたり曇ったりしている。満		
<u>おむね曇天となれり、琉球および</u>	<u>州と朝鮮はおおむね曇り、琉球(沖縄)</u>		
<u> 台湾にてはこれ反して快晴なり。</u>	<u>と台湾では、反して快晴。。</u>		

B.2.2 Semantic textual similarity on weather report commentary using SentenceBERT (SBERT)

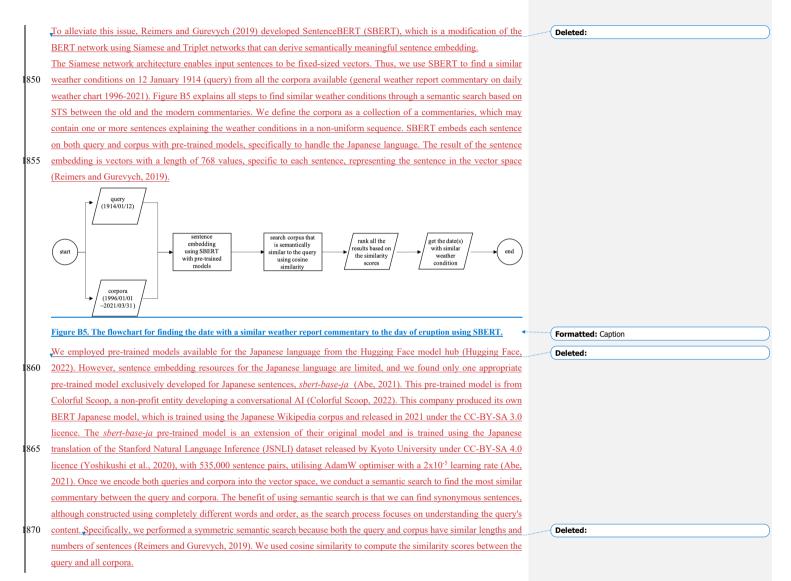
To find similar weather, both commentaries should have an equivalent meaning that is semantically similar, regardless of the sentence structures and words used to form the sentence. The vast improvements in machine learning (ML) and artificial intelligence (AI), especially in NLP tasks, allow us to find the similarity between two sentences using semantic textual similarity (STS) by embedding all sentences first. Word embedding represents words in the form of a real-valued vectors. This process encodes the meaning of a word such that the words that are closer in the vector space will be similar in meaning (Jurafsky and James, 2000). Sentence embedding is simply the sum of the individual word embedding (Cer et al., 2017). STS scores semantic similarity in varying degrees (e.g., a vehicle and a car are more similar than a wave and a car) instead of binary (e.g., a vehicle and a car are the same, and a wave and a car are not the same). STS provides a unified framework that allows for the extrinsic evaluation of multiple semantic components. This framework includes word sense disambiguation and induction, lexical substitution, semantic role labelling, multi-word expression detection and handling, anaphora and correference resolution, time and date resolution, named-entity handling, under-specification, hedging, semantic scoping and discourse analysis (Agirre et al., 2012).

Here, we can consider a pair of sentences to be similar (in this case, the weather condition) if they are located close to each other, as both sentences are mapped to a dense vector space (Reimers and Gurevych, 2019, 2020). Google has led the current state-of-the-art NLP development with its technology, Bidirectional Encoder Representations from Transformers (BERT), which can handle various NLP tasks accurately, including question answering, sentence classification, and sentence-pair

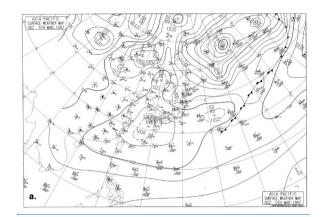
810 regression (Devlin et al., 2019). However, BERT is not optimal for handling several tasks, such as large-scale semantic similarity comparison, clustering, and information retrieval via semantic search (Reimers and Gurevych, 2020). Moved up [17]: This commentary agrees with the meteorological report provided in the chronicles (Omori, 1914). Therefore, we omitted the last phrases and focused only on the weather on the main islands of Japan. Then, further converted it to a form similar to the commentary in the current daily weather charts. The new phrases in the last column give more brief comments that match with the contemporary style in the modern weather charts while still maintaining the important features (high-pressure all over Japan, clear/sunny weather in main islands, Kyushu is cloudy, and Hokkaido is either clear or cloudy with strong south-westerly winds). Then, we compared these

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This commentary agrees with the meteorological report provided in the chronicles (Omori, 1914). Therefore, we omitted the last phrases and focused only on the weather on the main islands of Japan. Then, further converted it to a form similar to the commentary in the current daily weather charts. The new phrases in the last column give more brief comments that match with the contemporary style in the modern weather charts while still maintaining the important features (high-pressure all over Japan, clear/sunny weather in main islands, Kyushu is cloudy, and Hokkaido is either clear or cloudy with strong south-westerly winds). Then, we compared these phrases in contemporary Jananese with all the commentaries found in the modern weather chart (from 1 January 1996 to 31 March 2021) provided by the JMA in winter (from January 1996 to March 2021), from the historical archive collected by the National Institute of Informatics (Database of Weather Charts for Hundred Years, 2022). These commentaries explain the weather conditions on a particular day in a random sequence, that is, not in a similar order to the processed commentary phrases (from 12 January 1914). Finally, we utilised the natural language processing (NLP) method to find semantic textual similarity between all commentaries with sentence embedding using SentenceBERT (Reimers and Gurevych, 2019)...



se	entences are	likely to be lo	cated next to	each other in the vecto	or space (Cer	et al., 2017; Reimers and Gurevych, 2019;		
	arsov et al., 20					- un, 2017, Itemete und		
_						(5)		
		$\frac{v(q) \cdot v(c)}{\ v(q)\ _2 \ v(c)\ _2}$				(5)		
			× 2			uted by the dot product and the magnitude		
						vector $v(c)$, respectively. Here, q is each		
se	ntence in the	e queries, which	h is the featur	e we want to find, and e	c is each sente	ence in all the commentaries collected from		
	996 to 2021.	•				J		Deleted: The results of the comparison are between {0} for the most dissimilar sentences, and {1} if a pair of sentences is very
Ta	<u>able B2. The t</u>	top 10 weather s	<u>imilarities sco</u> r	ores of the selected dates.				similar (Singhal, 2001). The last step is to find similar weather conditions by ranking all corpora based on their similarity scores with
			_	Date	STS Score	•		conditions by ranking all corpora based on their similarity scores with the query. In brief, we selected only the top highest result from the top 10 results obtained from the search process, as shown in Table
			-	<u>5 March 1997</u>	<u>0.9635</u>	-	\mathbf{N}	B2
			_	24 February 2006	<u>0.9575</u>			(Formatted Table
			_	<u>2 February 2011</u>	<u>0.9551</u>			
			-	<u>11 December 1996</u>	<u>0.9532</u>	-		
			_	<u>16 March 2001</u>	<u>0.9514</u>			
				<u>30 March 2009</u>	<u>0.9508</u>			
				<u>28 January 2009</u>	<u>0.9507</u>			
			-	29 December 1996	<u>0.9503</u>			
				<u>19 January 2019</u>	<u>0.9497</u>	_		
				<u>14 January 2005</u>	<u>0.9496</u>			
	_							
Ta	able B3: The t	top result of the	semantic searc	rch process with its weathe	er description f	om the commentary	······································	Moved down [19]: Table B3 describes the general commentary on the top results of the semantic search. The
	Date	STS Score		<u>Commentary (ja)</u>		<u>Commentary (en)</u>		commentary in English is a direct translation using DeepL translation software (DeepL Translator, 2022), with sentences in
						5~10°C above normal nationwide. A mobile	$\langle \rangle \rangle$	bold indicating weather conditions similar to the Taisho eruption
			<u>全国的に当</u>	平年より5~10℃高い。	日本付近	high-pressure system will cover the area		Deleted:
			<u>は移動性</u>	高気圧に覆われるが、	<u>大陸から</u>	around Japan, but a trough of pressure will		Formatted: Caption Moved (insertion) [18]
<u>19</u>	<u>997/03/05</u>	<u>0.963</u>	<u>気圧の谷</u> ;	が進んでくる。北海道	<u>iで曇りの</u>	move in from the continent. Mostly cloudy	V	Formatted Table
			<u>他は似中</u>	晴れ。夜になると日本	<i>、海側は曇</i>	in Hokkaido and sunny throughout.		
			<i>b</i> .	山陰や九州北部では雨	<u></u> ज्ञ	Cloudy over the Sea of Japan at night. Rain		
				<u></u>		in the San-in region and northern Kyushu.		Deleted: Table B3: The top result of the semantic search process with its weather description from the commentary
-							francis	Formatted: Caption



905 Figure B6. Surface weather chart for Asia Pacific region at 00 UTC 5 March 1997. Weather chart image is obtained from the JMA4 Weather Charts for Hundred Years, 2022; Daily Weather Chart, 2022). (Database of The results of the comparison are between {0} for the most dissimilar sentences, and {1} if a pair of sentences is very similar (Singhal, 2001). The last step is to find similar weather conditions by ranking all corpora based on their similarity scores

with the query. In brief, we selected only the top highest result from the top 10 results obtained from the search process, as
 shown in Table B2. Table B3 describes the general commentary on the top results of the semantic search. The commentary in English is a direct translation using DeepL translation software (DeepL Translator, 2022), with sentences in bold indicating weather conditions similar to the Taisho eruption.

B.2.3 Weather condition on the selected date

The date with the highest STS scores (5 March 1997) agreed well with the weather features mentioned in the chronicles. The weather on selected date was generally clear and sunny because of the high-pressure covering Japan's main islands. With the exception of some rain that the chronicles did not mention. This difference indicate that we may not observe the exactly similar weather completely. This slight difference may affect the ashfall distribution produced by the simulation slightly. Figure B6 depicts the weather chart for the chosen date, illustrating more apparent features analogous to the weather chart published during the eruption. Both the chosen date and the eruption date exhibit the same weather characteristics, as highpressure covering entire Japan regions, added with an influence of a low-pressure centre in the north. On the selected date, the high-pressure centre moved from the south and then covered Japan at night. The weather charts presented in Fig. B6 further substantiate the continuous format transition of surface weather analysis in Japan. The general weather report Formatted: Justified

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commentary, accompanied by the establishment of a daily weather chart, is proven to be critical information that clarifies the 925 weather conditions on the respective date. Another feature, not mentioned in the commentary was the wind conditions during the selected dates. The chronicles briefly Deleted: revealed that the wind was weak, making it possible to evacuate using boats from the island to escape the eruption (Omori, 1914; Kitagawa, 2015). To further analyse the weather similarities between the selected date and 12 January 1914, we drew Deleted: the condition of the wind using a wind rose graph as depicted in Fig. B7. h а 85.7 51.4 68.6 Wind Speed (m/s) 1930 Figure B7. The wind rose diagrams show the six-hourly wind direction on 5 March to 10 March 1997, corresponding to which∢ Formatted: Caption from the winds blow at (a) 1000 hPa altitude (~100 m asl.), and (b) 250 hPa altitude (~10 km asl.). The graph consists of winds observations for every 6 h from the start of the simulation (00 UTC on the selected date) until Deleted: the simulation process finished (00 UTC four days after the eruption start). We chose the winds observations both at the 1000 hPa pressure level (equivalent to the surface pressure) and 250 hPa pressure level (equivalent to the location of the 935 westerlies) based on the findings explained in Sect. 2. The surface winds characteristics differed in terms of directions but had a comparable velocity between the two selected dates. The winds surface winds during the eruption day are akin to the selected date, regarding the dominant southwesterly winds portrayed in Fig. B7. The upper winds in March 1997 had a stronger velocity but mostly came from the west, due to the westerlies cycle. 940 Author contributions Deleted: Contributions HR performed most of the work to produce the dataset by performing the data processing and simulations, analysing the Deleted: did Deleted: the results, and writing this manuscript. In addition, HT conceptualised the project and supervised both the simulations and the Deleted: doing results analysis. MI produced the Taisho eruption data used as the eruption source parameters used in the model to generate the dataset. HLT is the creator of the PUFF model, provided the original FORTRAN codes of the model and the default 1945 values of the parameters used in the simulation. TT verified the weather similarity process by incorporating the general Deleted:



weather report commentary inside the weather chart. SR assisted in developing a semantic search methodology to find the semantic similarities through the commentary.

Competing Interest

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The authors declare that they have no conflict of interest.

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