

The consolidated European synthesis of CO₂ emissions and removals for 1 EU27 and UK: 1990-2020 2

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61 62 Abstract

63 Quantification of land surface-atmosphere fluxes of carbon dioxide (CO₂) fluxes and their trends and 64 uncertainties is essential for monitoring progress of the EU27+UK bloc as it strives to meet ambitious targets 65 determined by both international agreements and internal regulation. This study provides a consolidated synthesis of 66 fossil sources (CO₂ fossil) and natural sources and sinks over land (CO₂ land) using bottom-up (BU) and top-down 67 (TD) approaches for the European Union and United Kingdom (EU27+UK), updating earlier syntheses (Petrescu et 68 al., 2020, 2021b). Given the wide scope of the work and the variety of approaches involved, this study aims to answer 69 essential questions identified in the previous syntheses and understand the differences between datasets, particularly 70 for poorly characterized fluxes from managed ecosystems. The work integrates updated emission inventory data, 71 process-based model results, data-driven sectoral model results, and inverse modeling estimates, extending the 72 previous period 1990-2018 to the year 2020 to the extent possible. BU and TD products are compared with European 73 National Greenhouse Gas Inventories (NGHGIs) reported by Parties including the year 2019 under the United Nations 74 Framework Convention on Climate Change (UNFCCC). The uncertainties of the EU27+UK NGHGI were evaluated 75 using the standard deviation reported by the EU Member States following the guidelines of the Intergovernmental 76 Panel on Climate Change (IPCC) and harmonized by gap-filling procedures. Variation in estimates produced with 77 other methods, such as atmospheric inversion models (TD) or spatially disaggregated inventory datasets (BU), 78 originate from within-model uncertainty related to parameterization as well as structural differences between models. 79 By comparing NGHGIs with other approaches, key sources of differences between estimates arise primarily in 80 activities. System boundaries and emission categories create differences in CO2 fossil datasets, while different land 81 use definitions for reporting emissions from Land Use, Land Use Change and Forestry (LULUCF) activities result in 82 differences for CO₂ land. The latter has important consequences for atmospheric inversions, leading to inversions 83 reporting stronger sinks in vegetation and soils than are reported by the NGHGI.

84 For CO₂ fossil emissions, after harmonizing estimates based on common activities and selecting the most 85 recent year available for all datasets, the UNFCCC NGHGI for the EU27+UK accounts for 3392 ± 49 Tg CO₂ yr⁻¹ 86 $(926 \pm 13 \text{ Tg C yr}^{-1})$, while eight other BU sources report a mean value of 3340 [3238,3401] [25th,75th percentile] Tg CO2 yr⁻¹ (948 [937,961] Tg C yr⁻¹). The sole top-down inversion of fossil emissions currently available accounts for 87 88 3800 Tg CO₂ yr⁻¹ (1038 Tg C yr⁻¹), a value close to that of the NGHGI, but for which uncertainty estimates are not 89 yet available. For the net CO₂ land fluxes, during the most recent five-year period including the NGHGI estimates, 90 the NGHGI accounted for -91 ± 32 Tg C yr⁻¹ while six other BU approaches reported a mean sink of -62 [-117,-49] 91 Tg C yr⁻¹ and a 15-member ensemble of dynamic global vegetation models (DGVMs) reported -69 [-152,-5] Tg C yr⁻¹ 92 ¹. The five-year mean of three TD regional ensembles combined with one non-ensemble inversion of -73 Tg C yr^{-1} 93 has a slightly smaller spread (0th-100th percentile of [-135,45] Tg C yr⁻¹), and was calculated after removing land-94 atmosphere CO₂ fluxes caused by lateral transport of carbon (crops, wood trade and inland waters) resulting in 95 increased agreement with the the NGHGI and bottom-up approaches. Results at the sub-sector level (Forestland, 96 Cropland, Grassland) show generally good agreement between the NGHGI and sub-sector-specific models, but results 97 for a DGVM are mixed. Overall, for both CO2 fossil and net CO2 land fluxes, we find current independent approaches





98are consistent with the NGHGI at the scale of the EU27+UK. We conclude that CO_2 emissions from fossil sources99have decreased over the past 30 years in the EU27+UK, while large uncertainties on net uptake of CO_2 by the land100surface prevent trend identification. In addition, a gap on the order of 1000 Tg C yr⁻¹ between CO_2 fossil emissions101and net CO_2 uptake by the land exists regardless of the type of approach (NGHGI, TD, BU), falling well outside all102available estimates of uncertainties. However, uncertainties in top-down approaches to estimate CO_2 fossil emissions103remain uncharacterized and are likely substantial. The data used to plot the figures are available at104https://doi.org/10.5281/zenodo.7365863.

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106 1. Introduction

107 Atmospheric concentrations of greenhouse gasses (GHGs) reflect a balance between emissions from both 108 human activities and natural sources, and removals by the terrestrial biosphere, oceans, and atmospheric oxidation. 109 Increasing levels of GHG in the atmosphere due to human activities have been the major driver of climate change since the pre-industrial period (IPCC, 2021). In 2020, GHG mole fractions reached record highs, with globally 110 111 averaged mole fractions of 413.2 parts per million (ppm) for carbon dioxide (CO₂), representing 149% of the preindustrial level (WMO, 2021). The rise in CO₂ concentrations in recent decades is caused primarily by CO₂ emissions 112 from fossil sources. Globally, fossil emissions in 2020 (excluding the cement carbonation sink) totalled 9.5 ± 0.5 Gt 113 114 C yr⁻¹ (34.8 \pm 1.8 Gt CO₂ yr⁻¹), with expectations to rise in 2021 as the world recovered from the first year of the 115 Covid-19 pandemic (Friedlingstein et al., 2022). In contrast, global net CO₂ emissions from land use and land use 116 change (LULUC, primarily deforestation) estimated from bookkeeping models and dynamic global vegetation models 117 (DGVMs) were estimated to have a small decreasing trend over the past two decades, albeit with low confidence 118 (Friedlingstein et al., 2022). This decrease, however, is almost an order of magnitude less than the growth in fossil 119 emissions over the same period, and therefore the total fossil and net LULUC flux has still increased.

As all countries in the EU27+UK are Annex I Parties1 to the United Nations Framework Convention on 120 121 Climate Change (UNFCCC), they prepare and report national GHG emission inventories (NGHGIs) on an annual basis. These inventories contain annual timeseries of each country's GHG emissions from the 1990 base year² until 122 123 two years before the year of reporting and were originally set to track progress towards their reduction targets under the Kyoto Protocol (UNFCCC, 1997). Annex I NGHGIs are reported according to the Decision 24/CP.19 of the 124 125 UNFCCC Conference of the Parties (COP) which states that the national inventories shall be compiled using the 126 methodologies provided in the IPCC Guidelines for National Greenhouse Gas Inventories (IPCC, 2006). The 2006 127 IPCC Guidelines provide methodological guidance for estimating emissions for well-defined sectors using national 128 activity and available emission factors. Decision trees indicate the appropriate level of methodological sophistication

¹ Annex I Parties include the industrialized countries that were members of the OECD (Organization for Economic Co-operation and Development) in 1992 plus countries with economies in transition (the EIT Parties), including the Russian Federation, the Baltic States, and several central and eastern European states (UNFCCC, https://unfccc.int/parties-observers, last access: February 2022).

² For most Annex I Parties, the historical base year is 1990. However, parties included in Annex I with an economy in transition during the early 1990s (EIT Parties) were allowed to choose one year up to a few years before 1990 as reference because of a non-representative collapse during the breakup of the Soviet Union. For the EU27+UK, this includes Bulgaria (1988), Hungary (1985–1987), Poland (1988), Romania (1989), and Slovenia (1986).





129 (Tiered methods) based on the absolute contribution of the sector to the national GHG balance and the country's 130 national circumstances (availability and resolution of national activity data and emission factors). Generally, Tier 1 131 methods are based on global or regional default emission factors that can be used with aggregated activity data, while 132 Tier 2 methods rely on country-specific factors and/or activity data at a higher category resolution. Tier 3 methods are 133 based on more detailed process-level modeling or in some cases facility-level emission observations. Annex I Parties 134 are furthermore required to estimate and report uncertainties in emissions (95 % confidence interval) following the 2006 IPCC guidelines using, as a minimum requirement, the Gaussian error propagation method (approach 1). Annex 135 136 I Parties are furthermore encouraged to use Monte-Carlo methods (approach 2) or a hybrid approach. Additional information on the NGHGIs can be found in Appendix A1. 137

138 In addition to the NGHGIs, other research groups and international institutions produce independent estimates of national GHG emissions with two approaches: atmospheric inversions (top-down, TD) and GHG 139 140 inventories based on the same principle as NGHGIs but using slightly different methods (tiers), activity data, and/or emissions factors (bottom-up, BU). The current work has a strong focus on the EU27, and therefore sits within the 141 142 context of recent legislation passed by the European Parliament concerning commitments for the LULUCF sector to 143 achieve the objectives of the Paris Agreement and the reduction target for the Union (EU, 2018a and the proposed 144 amendments, EU, 2021a). This legislation requires that, "Member States shall ensure that their accounts and other 145 data provided under this Regulation are accurate, complete, consistent, comparable and transparent". The TD and BU 146 methods discussed below include the most up-to-date publicly available spatially explicit information, which can help 147 provide a quality check and increase public confidence in NGHGIs.

148 The work presented in this paper covers dozens of distinct datasets and models, in addition to the individual 149 country submissions to the UNFCCC of the EU Member States and the UK. As Annex I Parties, the NGHGIs of the 150 EU Member States and the UK are consistent with the general guidance laid out in IPCC (2006) yet still differ in 151 specific approaches, models, and parameters, in addition to definitional differences in the underlying system 152 boundaries and activity datasets. A comprehensive investigation of detailed differences between all datasets is beyond 153 the scope of this paper, though systematic analyses have been previously made for specific sectors (e.g. AFOLU³ -154 Petrescu et al., 2020; previous synthesis to this work - Petrescu et al., 2021b; FAOSTAT versus UNFCCC NGHGIs -155 Tubiello et al., 2021, Grassi et al., 2022a; UNFCCC versus bookkeeping models - Grassi et al, 2022b; and UNFCCC 156 versus inversions - Deng et al., 2021) and by the Global Carbon Project CO₂ syntheses (e.g., Friedlingstein et al., 2022). Every year (time "t") the Global Carbon Project (GCP) in its Global Carbon Budget (GCB) quantifies large-157 158 scale CO₂ budgets up to the previous year ("t-1"), bringing in information from global to large latitude bands, including 159 various observation-based flux estimates from BU and TD approaches (Friedlingstein et al., 2022). The current 160 manuscript, given the focus on a single region ("Europe") with extensive data coverage, dives into more detail than the GCB, including sector-specific models related to LULUCF (e.g., Forest land, Grassland, Cropland) and making 161 heavy use of the EU27+UK NGHGI in an effort to build mutual trust in the various approaches. Compared to Petrescu 162 163 et al. (2021b), the current work updates datasets, methods, and uncertainties.

³ We refer here to AFOLU as defined by the IPCC AR5: Agriculture, Forestry and Other Land Use.





164 BU observation-based approaches used in the GCB rely heavily on statistical data combined with Tier 1 and 165 Tier 2 approaches. In the current work, focusing on a region that is well-covered with data and models (EU27+UK), 166 BU also refers to Tier 3 process-based models (see Sect. 2). At regional and country scales, systematic and regular 167 comparison of these observation-based CO2 flux estimates with reported fluxes under the UNFCCC is more difficult. 168 Continuing our previous efforts within the European project VERIFY (VERIFY, 2022), the current study compares 169 observation-based flux estimates of BU versus TD approaches and compares them with NGHGI for the EU27-UK bloc and five sub-regions. VERIFY also provides, as a first attempt, similar comparisons for all European countries 170 171 (VERIFY Synthesis Plots, 2022). The methodological and scientific challenges to compare these different estimates 172 have been partly investigated before (Pongratz et al., 2021, Grassi et al., 2018a, for LULUCF; Andrew, 2020, for 173 fossil sectors) but such comparisons were not done in a systematic and comprehensive way, including both fossil and land-based CO₂ fluxes, before Petrescu et al. (2021b). 174

175 As Petrescu et al. (2021b) is the most comprehensive comparison of the NGHGI and research datasets 176 (including both TD and BU approaches) for the EU27+UK to date, the focus of the current paper is on improvement 177 of estimates in the most recent version in comparison with the previous one, including changes in the uncertainty 178 estimates and identification of the knowledge gaps and added value for policy making. Official NGHGI emissions are 179 compared with research datasets, including necessary harmonization of the latter on total emissions to ensure 180 consistency. Differences and inconsistencies between emission estimates were analyzed, and recommendations were 181 made towards future evaluation of NGHGI data. It is important to remember that, while NGHGIs include uncertainty 182 estimates, the "uncertainty analysis should be seen, first and foremost, as a means to help prioritize national efforts to 183 reduce the uncertainty of inventories in the future, and guide decisions on methodological choice" (Volume 1, Chapter 184 3, IPCC, 2006) and were therefore not developed to enable comparisons between countries or other datasets. In 185 addition, individual spatially disaggregated research emission datasets often lack quantification of uncertainty. Here, 186 we focus on the mean value and various percentiles (0th, 25th, 75th, 100th) of different research products of the same 187 type to get a first estimate of uncertainty (see Sect. 2). Not all models/inventories provided an update for v2021, and, 188 therefore, for the non-updated datasets the previously published timeseries are shown.

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190 2. CO₂ data sources and estimation approaches

191 The CO₂ emissions and removals in the EU27+UK estimated by inversions and anthropogenic emission 192 inventories resolved at the source category level were analyzed. At the time of this work, data of CO₂ fossil emissions 193 and CO₂ land⁴ emissions and removals (Tables 1 and 2) covered the period from 1990 to 2020, with some of the data

194 only available for shorter time periods. Since then, some datasets have been updated to include 2021, but not all, and

⁴ The IPCC Good Practice Guidance (GPG) for Land Use, Land Use Change and Forestry (IPCC, 2003) describes a uniform structure for reporting emissions and removals of greenhouse gasses. This format for reporting can be seen as "land based": all land in the country must be identified as having remained in one of six classes since a previous survey, or as having changed to a different (identified) class in that period. According to IPCC SRCCL: Land covers the terrestrial portion of the biosphere that comprises the natural resources (soil, near surface air, vegetation and other biota, and water) the ecological processes, topography, and human settlements and infrastructure that operate within that system". Some communities prefer "biogenic" to describe these fluxes, while others find this confusing as fluxes from unmanaged forests, for example, are "biogenic" but not included in inventories reported to the UNFCCC. As this comparison is central to our work, we decided that "land" as defined by the IPCC was a good compromise.





195 we made the decision to stay with the original time window for simplicity. The estimates are available both from 196 peer-reviewed literature and from new research results from the VERIFY project. BU results are compared to NGHGI 197 reported in 2021 (which contain the timeseries for 1990-2019). Data sources are summarized in Tables 1 and 2 with 198 the detailed description of all products provided in Appendices A1-A2. In Appendix A, the harmonized methodology 199 for calculation of uncertainties submitted by Member States to the UNFCCC in their National Inventory Reports 200 (NIRs) is explained. This includes the same 95 % confidence interval as is typically reported, but involved an 201 extensive gap-filling to cover more categories and more years than available in Petrescu et al. (2021b), which limited 202 uncertainty estimation to a single year.

203 BU anthropogenic CO2 fossil estimates include global inventory datasets such as the Emissions Database for 204 Global Atmospheric Research (EDGAR v6.0.), Statistical Review of World Energy by BP, the Carbon Dioxide 205 Information Analysis Center (CDIAC), the Global Carbon Project (GCP), the Energy Information Administration's 206 (EIA) "International" dataset, and the International Energy Agency (IEA) (see Table 1). These datasets are all 207 described in detail by Andrew (2020). CO2 land emission estimates are derived from BU biogeochemical models (e.g. 208 DGVMs, bookkeeping models, see Table 2). TD approaches include both high spatial resolution regional inversions 209 (CarboScopeReg, EUROCOM (Monteil et al., 2020), inversions based on the CIF-CHIMERE system (Berchet et al., 210 2021) and LUMIA) and coarser spatial resolution global inversions (GCP 2021: Friedlingstein et al., 2022). Most of 211 the inversions were carried out for CO₂ land emissions, with only a single inversion for CO₂ fossil emissions (CIF-212 CHIMERE). Note that CIF-CHIMERE provides estimates for both CO₂ land and CO₂ fossil from separate simulations. 213 These estimates are described in Sect. 2.3.

214 The sign of the fluxes is defined from an atmospheric perspective: positive values represent a net source to 215 the atmosphere and negative values a net removal from the atmosphere. As an overview of potential uncertainty 216 sources, Table B1 presents the use of emission factor data (EF), activity data (AD), and, whenever available, 217 uncertainty methods used for all CO2 land data sources in this study, in addition to more details on each model in 218 Appendices A. The referenced data used for the figures' replicability purposes are available for download (McGrath 219 et. al, 2022). Upon request, the codes necessary to plot the figures in the same style and layout can be provided. The 220 focus is on EU27+UK emissions. In the VERIFY project, an additional web tool was developed which allows for the 221 selection and display of all plots shown in this paper, not only for the EU Member States and UK but for a total of 79 222 countries and groups of countries in Europe (Table A1, Appendix A). The data is free and can be accessed upon 223 registration (VERIFY Synthesis Plots, 2022).

224 For the sake of harmonization, we report the mean values of all ensembles. For small sample sizes (e.g., the 225 regional inversions of CSR with four members), the literature does not give a clear indication on whether the mean or 226 the median is preferred; a preference for one or the other depends on what one wishes to demonstrate. In particular, 227 the median downplays the skewness of the data (outliers). We have taken efforts to exclude outliers from the datasets 228 used to construct ensembles, and consequently the datasets which remain should be randomly distributed. For this 229 reason, we display the mean for all ensembles. As the number of datasets in some ensembles is small (less than five), we display the minimum and maximum annual values for every year (i.e., the 0th/100th percentiles) to give an idea 230 231 of the spread. For ensembles with more than ten members (i.e., TRENDY), we show the mean and the 0th/100th





percentiles along with the 25th/75th percentiles in the figures. This combination demonstrates "more likely" and "possible" behavior; as only one ensemble has both bars, displaying them does not overwhelm the reader much more than the standard graphs, and we find the added information to be worth the trade-off. In the text, we report the mean and 0/100th percentiles for small ensembles and mean along with the 25th/75th for larger ensembles.

The current work extends Petrescu et al. (2021b) by updating the included datasets (both increasing the number of years covered and in some cases updating the model versions), adding datasets, and highlighting changes in terms of mean annual emissions and trends. For clarity, the data from Petrescu et al. (2021b) is labeled as v2019, while the latest results are labeled v2021.

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241 2.1. CO₂ anthropogenic emissions from NGHGI

242 The UNFCCC NGHGI (2021) estimates for the period 1990 to year t-2 (2019), collected for the EU27 and 243 UK, are the basis for this dataset. For historical reasons, a few EU countries provide data for a different base year than 1990⁵, yet it should be noted that regardless of the base year all countries of the EU27+UK bloc are obliged to report 244 estimates for the period 1990 to year t-2. The Annex I Parties to the UNFCCC are required to report annual GHG 245 246 inventories that include a NIR, with qualitative information on data and methods and a Common Reporting Format (CRF) set of tables that provide quantitative information on GHG emission by category. This annually updated dataset 247 248 includes anthropogenic emissions and removals. For the land-based sector, the land management proxy is used as a 249 way to report only anthropogenic fluxes (Grassi et al., 2018a, 2021). This proxy allows Member States to report all 250 fluxes coming from land designed as "managed" without trying to disentangle their natural and anthropogenic origins. Figure B1 shows the annual NGHGI (2021) anthropogenic CO2 timeseries disaggregated by sector in order to provide 251 252 context.

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254 2.2. CO₂ fossil emissions

255 CO2 fossil emissions occur when fossil carbon compounds are broken down via combustion or other non-256 combustive industrial processes. Most of these fossil compounds are in the form of fossil fuels, such as coal, oil, and 257 natural gas. Another source category of fossil CO₂ emissions is fossil carbonates, such as calcium carbonate and 258 magnesium carbonate, which are used in industrial processes. Because CO2 fossil emissions are largely connected 259 with energy, which is a closely tracked commodity group of high economic importance, there is a wealth of underlying 260 data that can be used for estimating emissions. However, differences in collection, treatment, interpretation and 261 inclusion of various factors - such as carbon contents and fractions of the fuel's carbon that is oxidized - lead to 262 methodological differences (Appendix A) resulting in differences of emissions between datasets (Andrew, 2020). 263 Atmospheric inversions for emissions of fossil CO₂ are not as established as their bottom-up counterparts (Brophy et

⁵ For most Annex I Parties, the historical base year is 1990. However, parties included in Annex I with an economy in transition during the early 1990s (EIT Parties) were allowed to choose one year up to a few years before 1990 as reference because of a non-representative collapse during the breakup of the Soviet Union (e.g., Bulgaria, 1988, Hungary, 1985–1987, Poland, 1988, Romania, 1989, and Slovenia, 1986).





264 al., 2019). The main reason is that the types of atmospheric monitoring instruments suitable for fossil CO₂ atmospheric 265 inversions have not yet been widely deployed (Ciais et al., 2015). One of the rare inversions is presented below. 266 In this analysis, the inventory-based bottom-up CO₂ fossil emissions estimates are separated and presented 267 per fuel type and reported for the last year when all data products are available (2017). This updates Andrew (2020) 268 and Petrescu et al. (2021b) which both report the year 2014. In order to provide a quasi-independent estimate of fossil 269 emissions assimilating satellite observations of the atmosphere, the CIF-CHIMERE model was used to produce a 270 fossil fuel CO2 emission estimate for the year 2017. CIF-CHIMERE is a coupling between the variational mode of 271 the Community Inversion Framework (CIF) platform developed in the VERIFY project (Berchet et al., 2021), the 272 CHIMERE chemistry transport model (Menut et al., 2013) and the adjoint of this model (Fortems-Cheiney et al., 273 2021a). To overcome the lack of CO₂ observation networks suitable for the monitoring of fossil fuel CO₂ emissions 274 at national scale, this inversion is based on the assimilation of satellite NO2 data, as NO2 is co-emitted with CO2 during 275 fossil fuel combustion. Recent top-down inversions of anthropogenic CO₂ emissions from Europe indicate that 276 uncertainties using satellite measurements of co-emitted NO2 are much lower than for co-emitted CO when deriving 277 fossil CO2 emissions (Konovalov et al., 2016). Therefore, results shown below only incorporate NO2 and not CO 278 observations. While the spatial and temporal coverage of the NO₂ observations is large, there are many factors that 279 determine the ratio of NO2 to CO2 emissions. Therefore, the influence of using NO2 observations in determining fossil 280 CO₂ emissions is subject to uncertainties which have not been characterized appropriately yet in the framework of 281 VERIFY. Here, this conversion relies heavily on the emission ratios per country, month and large sector of activity 282 from the TNO-GHGco-v3 inventory (Dellaert et al., 2021), which has been partly developed in VERIFY, and which 283 is based on the most recent UNECE-CLRTAP⁶ and UNFCCC official country reporting respectively for air pollutants 284 and greenhouse gasses. The detailed descriptions of each of the data products are found in Appendix A1.

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Table 1: Data sources for the anthropogenic CO₂ fossil emissions included in this study, all updated from Petrescu et

²⁸⁷ al. (2021b):

⁶ UNECE Convention on Long-Range Transboundary Air Pollution. https://unece.org/environment-policy/air





Anthropogenic CO ₂ fossil				
Data/model name	Contact / lab	Species / Period	Reference/Metadata	
UNFCCC NGHGI (2021)	UNFCCC	Anthropogenic fossil CO ₂ 1990-2019	IPCC (2006) UNFCCC NIRs/CRFs https://unfccc.int/ghg-inventories-annex-i- parties/2021 (UNFCCC, 2021a, 2021b) EDGAR v6 0	
Compliation of multiple CO ₂ fossil emission data sources (Andrew 2020) EDGAR v6.0, BP, EIA, CDIAC, IEA, GCP, CEDS, PRIMAP	CICERO	totals and split by fuel type 1990-2018 (or last available year)	 Z EDGAR v6.0 kttps://edgar.jrc.ec.europa.eu/ t BP 2021 report (BP, 2021) EIA https://www.eia.gov/beta/international/data/toowser/views/partials/sources.html CDIAC https://energy.appstate.edu/CDIAC (Gilfill and Marland, 2021) IEA : www.iea.org CEDS https://github.com/JGCRI/CEDS (O'Rourke al., 2021) GCP (Friedlingstein et al., 2022) PRIMAP-hist (Gütschow et al., 2021) https://doi.org/10.5281/zenodo.4479171 	
Fossil fuel CO ₂ inversions	LSCE	Inverse fossil fuel CO ₂ emissions 2005-2020	Fortems-Cheiney et al. (2021) Fortems-Cheiney and Broquet (2021)	

288 2.3. CO₂ land fluxes

Data products from BU and TD CO₂ land fluxes including CO₂ emissions and removals from land use, land use change, and forestry (LULUCF) activities are summarized in Table 2. All models and approaches produce an estimate of the net carbon flux from the land surface including uptake through photosynthesis and emission through respiration and/or disturbances. The details may vary significantly between approaches, however. Attempts are made where possible to harmonize input data and compare results which roughly correspond to similar categories included in the NGHGI. Further details are described throughout the rest of this article. As with CO₂ fossil fluxes, the primary distinctions are between the NGHGI, other bottom-up approaches, and top-down approaches. The situation becomes

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297 more complicated for CO₂ land fluxes due to the inclusion of approaches which only address a single land use class
 298 (e.g., Forest land).

299 For the analysis at category level, the CO₂ net emissions from the LULUCF sector that are primarily 300 considered in this synthesis are from three land use classes7 (Forest land, Cropland, and Grassland), each split into a 301 land class remaining in the same land class⁸ or a land class converted to another class. The NGHGIs are the only 302 results discussed here which make use of this transition period, but the distinction is important so as to inform which 303 NGHGI categories to use in the comparison. Wetlands, Settlements, Other land, and Harvested wood products (HWP) 304 categories are included in the discussion on total LULUCF activities in Sect. 3.3.1, 3.3.3 and 3.3.4. Not all the classes 305 reported to the UNFCCC are present in FAOSTAT or other models. Some models are sector-specific (e.g., Forest land) while other models include a larger subset of the six UNFCCC classes (e.g., DGVMs which simulate Forest 306 307 land, Grassland, and Cropland). The notations FL, CL and GL are used to indicate total emissions and removals from 308 the respective Forest land, Cropland and Grassland land use categories (i.e. the remaining + conversions to these classes). The notations "FL-FL", "CL-CL" and "GL-GL" are used to indicate emissions and removals from respective 309 310 forest, cropland and grassland areas which have remained in the same class from year to year, or in the case of NGHGI 311 lands that have not undergone conversion within the aforementioned transition period (e.g. t-20).

312 The results from sector-specific models reporting carbon fluxes for FL-FL (EFISCEN-Space and CBM), CL and GL (EPIC-IIASA and ECOSSE) are presented separately from the models and datasets including multiple land 313 314 use categories and simulating land use changes: FAOSTAT (version 2021), the DGVM ensemble TRENDY v10 (Friedlingstein et al., 2022; Le Quéré et al., 2009), the ORCHIDEE and CABLE-POP DGVMs forced by high 315 resolution meteorological data as part of the VERIFY project, and the two bookkeeping approaches of H&N 316 317 (Houghton & Nassikas, 2017) and BLUE (Hansis et al., 2015). BLUE includes two simulations with different landuse forcing, one made for the VERIFY H2020 project and one for the GCP 2021 (Friedlingstein et al., 2022). For CL 318 319 and GL both the EPIC-IIASA and ECOSSE sector-specific models reported updates, although ECOSSE only updated 320 results for GL. Process included in all the products are summarized in Appendix A2 and Table B2.

321 The two updated inverse model ensembles presented are the GCP2021 for the period 2010-2020 322 (Friedlingstein et al., 2022) and EUROCOM for the period 2009-2018 (Monteil et al., 2020; Thompson et al., 2020). 323 The GCP inversions are global and include CarbonTracker Europe (CTE: van der Laan-Luijkx et al., 2017), CAMS 324 (Chevallier et al., 2005), the Jena CarboScope (Rödenbeck, 2005), NIESMON-CO₂ (Niwa et al., 2017), CMS-Flux 325 (Liu et al., 2021) and UoE (Feng et al., 2016). The EUROCOM inversions are regional, with a domain limited to 326 Europe and higher spatial resolution atmospheric transport models, with four inversions covering the entire period 2009-2018 as analyzed in Thompson et al. (2020). All inversions provide Net Ecosystem Exchange (NEE) fluxes. 327 328 These inversions make use of more than 30 atmospheric observing stations within Europe, including flask data and

⁷ According to 2006 IPCC guidelines the LULUCF sector includes six management classes (Forest land, Cropland, Grassland, Wetlands, Settlements and Other land)

⁸According to 2006 IPCC guidelines, land converted to a new category should be reported in a "conversion" category for N years and then moved to a "remaining" category, unless a further change occurs. Converted land refers to CO₂ emissions from conversions to and from all six classes that occurred in the previous N years. By default, N is equal to 20, although the guidelines recognize that longer times may be necessary in temperate and boreal environments for the dead biomass and soil carbon pools to reach the new equilibrium. Member States have the freedom to select a length of time appropriate to their own circumstances.





329 continuous observations and work at typically higher spatial resolution than the global inversion models (Table 2). 330 The prior anthropogenic emissions provided for all regional inversions reported here (i.e., EUROCOM, EUROCOM 331 drought 2018, VERIFY CSR, VERIFY CIF-CHIMERE, and VERIFY LUMIA) are all based on EDGAR v4.3, BP 332 statistics, and TNO datasets by generating spatial and temporal distributions through the COFFEE approach 333 (Steinbach et al., 2011). Small differences exist between exact versions used by the different groups. The prior 334 anthropogenic emissions for the GCP global inversions, GridFEDv2021 and v2022, are also based on EDGARv4.3.2 335 (Janssens-Maenhout et al., 2019). Overall, differences in the prior anthropogenic emissions are not expected to explain 336 the large differences seen between the different regional biogenic inversions nor between the regional and global 337 biogenic inversions, but efforts should be continued to harmonize them to the greatest extent possible in future 338 intercomparisons.

339 Additional inversions for Europe from three regional scale inversion systems are analyzed. Two of these 340 systems are part of the EUROCOM ensemble, but new runs were carried out for the VERIFY project. The CarboScopeRegional (CSR) inversion system has performed additional runs for VERIFY for the years 2006-2020 341 342 with multiple ensemble members differing by biogenic prior fluxes and assimilated observations. The results are 343 plotted separately to illustrate two points: 1) that the CSR simulations for VERIFY are not identical to those submitted 344 to EUROCOM (VERIFY runs from CSR included several sites that started shortly before the end of the EURCOM 345 inversion period), and 2) the CSR model was used in four distinct runs in VERIFY. Note that the ensemble members 346 differ from previous years (the spatial correlation length is kept constant this year, while more prior fluxes are used). 347 By presenting CSR separate from the EUROCOM results, one can get an idea of the uncertainty due to various model 348 parameters in one inversion system with one single transport model. The LUMIA inversion system submitted four 349 simulation results to the VERIFY project, based on the 2018 Drought Task Force project (labeled here as EUROCOM, 350 Thompson et al., 2020). The primary difference is that the years 2019-2020 were added based on boundary conditions 351 using TM5 and ERA5 meteorological data. The four different variants include one reference simulation and three simulations which change spatial correlation lengths, the number of observation sites, and the magnitude of 352 353 uncertainties in the boundary conditions. As one of the variants is only available for 2019-2020 (changing the 354 uncertainties in the boundary conditions), this variant was dropped from the results and only the remaining three 355 simulations are presented, covering the period 2006-2020.

An inversion of the NEE over 2005-2020 from the CIF-CHIMERE variational inversion system is also analyzed. The configuration of this inversion is close to that of the PYVAR-CHIMERE NEE inversions in the EUROCOM ensembles and follows the general principles of Broquet et al. (2013). However, it uses distinct inputs, which play a critical role in the inversion, such as a more recent ORCHIDEE simulation as prior estimate of the NEE and a more recent CAMS global inversion to impose the regional CO₂ boundary conditions.

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Table 2: Data sources for the land CO₂ emissions included in this study. Details are found in Appendix A2.





	NGHGI net CO ₂ land flux				
Data source	Contact / lab	Variables Period (timestep) Resolution	References	Status compared to Petrescu et al. (2021b)	
UNFCCC NGHGI (2021)	Member State inventory agencies Annual uncertainty gap- filling for total LULUCF by Environment Agency Austria (EAA).	LULUCF Net CO ₂ emissions/removals. 1990-2019 (1Y) Country-level	IPCC (2006) UNFCCC CRFs https://unfccc.int/process-and- meetings/transparency-and- reporting/reporting-and- review-under-the- convention/greenhouse-gas- inventories-annex-i- parties/national-inventory- submissions-2019	Updated	
	Inventory and mo	odel estimates of net CO ₂	land flux		
ORCHIDEE	LSCE	CO ₂ fluxes from all ecosystems reported as Net Biome Productivity (NBP). 1990-2020 (3H) 0.125° x 0.125°	Ducoudré et al. (1993) Viovy et al. (1996) Polcher et al. (1998) Krinner et al. (2005)	Updated	
CABLE-POP	Western Sydney University	CO ₂ fluxes (NBP). Model includes N cycling. 1990-2020 (1M) 0.125° x 0.125°	Haverd et al. (2018)	New	
TRENDY v10	MetOffice UK	CO ₂ fluxes (NBP) 15 models (all except ISAM) 1990-2020 (3H-1M) 0.125° x 0.125°	Friedlingstein et al. (2022; Table 4)	Updated	





CO ₂ emissions from inland waters	ULB	Average C fluxes from rivers, lakes and reservoirs, with lateral C transfer from soils. 1990-2018 (-) 0.1° x 0.1°	Lauerwald et al. (2015) Hastie et al. (2019) Raymond et al. (2013)	Not updated
СВМ	EC-JRC	CO ₂ fluxes (NBP) as historical 2000-2015 and extrapolation for 2017-2020 (1Y) Country-level	Kurz et al. (2009) Pilli et al. (2022)	Updated
ECOSSE	UNIABDN	CO ₂ fluxes (NBP) from croplands and grassland ecosystems. Crops: 1990-2020 (1Y) Grass: 1990-2018 (1Y) 0.125° x 0.125°	Bradbury et al. (1993) Coleman (1996) Jenkinson (1977, 1987) Smith et al. (1996, 2010a,b)	Updates only for croplands
EFISCEN-Space	WUR	CO ₂ fluxes (NBP): single average value for 5 year periods, replicated on a yearly time axis. 0.125° x 0.125°	Verkerk et al. (2016) Schelhaas et al. (2017, 2020) Nabuurs et al. (2018)	Updates for 15 countries
EPIC-IIASA	IIASA	CO ₂ fluxes (NBP) from cropland 1991-2020 (1M) 0.125° x 0.125°	Balkovič et al. (2013, 2018, 2020) Izaurralde et al. (2006) Williams et al. (1990)	Updated for croplands, new estimates for grasslands
BLUE (VERIFY) and BLUE (GCP)	LMU Munich	CO ₂ fluxes from land use change. VERIFY: 1990-2019 (1Y) GCP: 1990-2020 (1Y) 0.25° x 0.25°	Hansis et al. (2015) Ganzenmüller et al. (2022) - VERIFY Friedlingstein et al. (2022) - GCP2021	Updated
H&N	Woodwell Climate Research Center	CO ₂ fluxes from land use change. 1990-2020 (1Y) Country-level	Houghton and Nassikas (2017)	Updated
FAO	FAOSTAT	CO ₂ emissions / removal from LULUCF processes. 1990-2020 (1Y) Country-level	FAO (2021) Federici et al. (2015) Tubiello et al. (2021)	Updated
CO2 atmospheric inversion estimates				





CSR inversions for VERIFY	MPI -Jena	Total CO ₂ inverse flux (NBP) 2006-2020 (3H) 0.5° x 0.5°	Kountouris et al. (2018 a,b)	Updated
LUMIA	Lund University (INES)	Total CO ₂ inverse flux (NBP) 2006-2020 (1W) 0.25° x 0.25°	Monteil and Scholze (2021)	New
CIF-CHIMERE	LSCE	Total CO ₂ inverse flux (NBP) 2005-2020 (3H) 0.5° x 0.5°	Berchet et al. (2021) Broquet et al. (2013)	New
GCP 2021 global inversions (CTE, CAMS, CarboScope, NISMON-CO ₂ , UoE, CMS- Flux)	GCP	Total CO ₂ inverse flux (NBP) Six inversions 2010-2020 (various)	Friedlingstein et al. (2022) Van der Laan-Luijk et al. (2017) Chevallier et al. (2005) Rödenbeck et al. (2005) Niwa et al. (2017) Feng et al. (2016) Liu et al. (2021)	Updated
EUROCOM regional inversions (CSR, LUMIA, PYVAR)	LSCE, ULUND, MPI-Jena, NILU	Total CO ₂ inverse flux (NBP) Three inversions 2009-2018 (3H-1M)	Monteil et al. (2020) Thompson et al. (2020)	Updated (also replaced CSR with the mean of the four runs submitted to VERIFY). FLEXINVERT and NAME are not included (Fig. A5)

363

All of the bottom-up models in this work require external forcing datasets. In the context of the VERIFY project (VERIFY, 2022), an effort was made to provide a single, harmonized version of several kinds of data (meteorological, land use/land cover, and nitrogen deposition) on a high-resolution grid over Europe. These datasets were then made available to all of the modeling groups to use in their simulations. Such a practice is common in model intercomparison projects. However, as the models in Table 2 are not all the same type, data harmonization presented more of a challenge in this work as not all models use the same inputs. All of the datasets described in Appendix A2 were used by at least one modeling group in this work.

371 3. Results and discussion

372 3.1. Overall NGHGI reported anthropogenic CO₂ fluxes

373 In 2019, the UNFCCC NGHGI (2021) net CO₂ flux estimates for EU27+UK, accounted for 3.01 Gt CO₂

from all sectors (including LULUCF) and 3.28 Gt CO₂ excluding LULUCF (Fig. B1), corresponding to a net sink of



375 LULUCF of -0.27 ± 0.11 Gt CO₂. In 2019, few large economies accounted for the majority of EU27+UK emissions, 376 with Germany, UK, Italy and France representing 53 % of the total CO₂ emissions (excluding LULUCF). For the 377 LULUCF sector, the countries reporting the largest CO₂ sinks in 2019 were Italy, Spain, Sweden, and France 378 accounting for 56 % of the overall EU27+UK sink. Only a few countries (Czech Republic, The Netherlands, Ireland 379 and Denmark) reported a net LULUCF source in 2019. Some countries, like Portugal, report sources in some years 380 due to wildfires, with sinks in other years. The NGHGI shows minimal inter-annual variability (largely due to 381 methodology), and consequently the 2019 values are indicative of longer-term averages, showing a constant trend 382 between 2017-2019.

CO₂ fossil emissions reported by Member States are dominated by the energy sector (energy combustion and
 fugitives) representing 92 % of the total EU27 + UK CO₂ emissions (excluding LULUCF) or 3.02 Gt CO₂ yr⁻¹ in 2019.
 The Industrial Process and Product Use (IPPU) sector contributes 7.6 % or 0.2 Gt CO₂ yr⁻¹. CO₂ emissions reported
 as part of the agriculture sector cover only liming and urea application, UNFCCC categories 3G and 3H⁹ respectively.
 Together with waste, in 2019 the emissions from agriculture represent 0.4 % of the total UNFCCC CO₂ emissions in
 the EU27+UK.

An overview of all CO₂ fossil and land datasets in this work (Fig. 1) leads to a series of conclusions: 1) Regardless of the method used (NGHGI, bottom-up models, top-down models), the timeseries of annual fluxes from fossil CO₂ emissions rest almost one order of magnitude higher than removals from CO₂ uptake/removal by the land surface and well outside uncertainty estimates; 2) Uncertainties are much larger in the LULUCF estimates than in the fossil CO₂ estimates (both for total LULUCF and for individual components of FL, CL, and GL); 3) Interannual variability (IAV) is much more present in non-NGHGI LULUCF datasets than in NGHGI LULUCF datasets or any of the fossil datasets.

The overall message that fossil CO₂ emissions exceed the land sink (Fig. 1a-c) is the same as found in the Global Carbon Budget (Friedlingstein et al., 2022), although the difference is larger in the EU27+UK. Contrary to the GCB, however, fossil CO₂ emissions in the EU27+UK have decreased over the past three decades. Again, this finding is supported by the NGHGI, bottom-up models, and a single atmospheric inversion. Similarly, carbon uptake by the land surface has remained more or less stable over the past three decades, with the vast majority of that occurring in forests. While the latter conclusion is clear in the NGHGI (Fig. 1d), very large spreads among bottom-up sectorial models lead to more uncertainty (bottom-center).

The difference in uncertainty between the estimates of fossil CO₂ emissions and CO₂ uptake/removal by the land surface is also striking. Eight bottom-up models produce a 25-75 % percentile which is almost invisible on the scale of the graph (center-top, gray shading). On the other hand, four models estimating Grassland emissions/removals produce an error bar that covers the bottom part of the graph and masks any apparent trend (bottom-center, light green shading). A similar conclusion can be drawn from top-down estimates of LULUCF fluxes (top-right, blue shading). Additional work on reducing the uncertainty of LULUCF fluxes in the EU27+UK is highly welcome.

409

⁹ 3G and 3H refer to UNFCCC category activities, as reported by the standardized common reporting format (CRF) tables, which contain CO₂ emissions from agricultural activities: liming and urea applications.







410

Figure 1: A synthesis of all the CO_2 net fluxes shown in the work for the EU27+UK. The estimates are divided by 411 approach: NGHGI estimates (panels a, d); bottom-up methods (b, e); and top-down methods (c). Panels (d) and (e) 412 413 include a breakdown of the LULUCF flux into three of the dominant components: FL, GL, and CL. Such a breakdown 414 is not provided for NHGHI CO2 fossil as partitioning of bottom-up CO2 fossil datasets corresponding to UNFCCC 415 NGHGI categories is not currently available. The NGHGI UNFCCC uncertainty is calculated for submission year 416 2021 as the relative error of the NGHGI value, computed with the 95 % confidence interval method gap-filled and 417 provided for every year of the timeseries, except for FL, GL, and CL which are taken directly from the EU NIR (2021). 418 Shaded areas for the other estimates represent the 0th-100th percentiles for groups with fewer than seven members, 419 and the 25th-75th percentile for groups with seven or more members. Ensembles (e.g., TRENDY v10) are included 420 in the above only their mean values, to avoid more heavily weighting the ensembles compared to the other datasets. 421

422 Several caveats remain with this overall synthesis. First, the timeseries were combined rather naively in Fig. 1 423 by taking the mean of annual timeseries for each dataset discussed below. This leads to, for example, the 15-member 424 TRENDY ensemble being given identical weight as the ORCHIDEE high-resolution simulation over Europe. This 425 was done to weigh more heavily the regional approaches under the assumption that higher resolution simulations and 426 more region-specific input data will lead to more accurate results. While the latter assumption appears reasonable, 427 the first assumption can be disputed. Second, only a single top-down result for fossil CO₂ emissions is currently available, preventing an estimate of the uncertainty for this approach. Third, sector models were combined 428 429 disregarding distinctions between those models estimating "Remain" and "Total" fluxes. These points are discussed





in more detail in the following sections. However, addressing these points is highly unlikely to alter the overallconclusions in this section.

432

433 3.2. CO₂ fossil emissions

434 The inventory-based fossil CO₂ estimates from nine data sources (and some subsets) are presented as timeseries 435 (1990-last available year) based on Andrew (2020) with the objective to explore differences between datasets and 436 visualize trends (Fig. 2). Because the emissions source coverage (also called the "system boundary") of datasets 437 varies, comparing total emissions from these datasets is not a like-for-like comparison. Therefore, some harmonization 438 of system boundaries prior to comparison is needed. This harmonization relies on specifying the system boundary of 439 each dataset and, where possible, removing emission sources to produce a near-common system boundary. For 440 example, if IEA doesn't include any carbonates, then carbonates are removed from all emissions datasets that report 441 these separately. UNFCCC (CRFs) Energy+IPPU, CDIAC, CEDS, PRIMAP, and GCP include Energy sector plus 442 all fossil fuels in IPPU; EIA, EDGAR and BP include some fossil fuels in IPPU, while EIA and BP include bunker 443 fuels as well. UNFCCC CRFs include Energy total and Energy combustion. Further details on how data sets are 444 harmonized are provided by Andrew (2020). Because of differing levels of detail provided by datasets, it isn't possible 445 to do this perfectly, but the approximate harmonization gives something closer to a like-for-like comparison, with the 446 legend in Fig. 2 indicating the most significant remaining differences. The pre-harmonization curves are shown in 447 Appendix A (Fig. A1) for reference.

448

449







451 Figure 2: Comparison of EU27+UK fossil CO₂ emissions from multiple inventory datasets with system boundaries
452 harmonized as much as possible. Harmonization is limited by the disaggregated information presented by each
453 dataset. CDIAC does not report emissions prior to 1992 for former-Soviet Union countries. CRF: UNFCCC NGHGI

454 from the Common Reporting Format tables. The pre-harmonization figure is shown in Fig. A1.

455 Given the remaining differences in system boundaries after harmonization, most datasets agree well 456 (Andrew, 2020). In response to inconsistencies identified in this work, the EIA recently corrected some double-457 counting of emissions from liquid fuels and has revised its estimates of total emissions down about 10 % for the 458 EU27+UK (pers. comm., US Energy Information Agency, February 2022). For comparison, applying a similar 459 harmonization procedure to the UNFCCC NGHGI and retaining only Fuel combustion (1A), Fugitive emissions (1B), 460 Chemical industry (2B), Metal industry (2C), Non-energy products from fuels and solvent use (2D), and Other (2H) results in emissions of 3392 ± 49 Tg CO₂ yr⁻¹ (926 ± 13 Tg C yr⁻¹) for the year 2017, where the uncertainty was 461 462 propagated through quadrature using the gap-filled uncertainties described in this work and taking the total sector 463 uncertainty if the category uncertainty was not available. This mean value falls within the 25th-75th percentiles of 464 the eight other harmonized BU sources ([3238,3401] Tg CO₂ yr⁻¹).

465 The sole available inversion for CO₂ fossil fluxes is produced by the CIF-CHIMERE model, shown in Fig. 466 1c and Fig. B3 (for a single year). The inversion yields plausible and consistent fossil emission estimates compared 467 to nine bottom-up estimates from BU datasets with global coverage (Fig. 1b,c,B3). Uncertainties of CIF-CHIMERE 468 inversion estimate have not yet been quantified, however they are likely largely driven by large uncertainties in the 469 input data. The satellite observations of NO₂ have large uncertainties, which partly explains the small departure from 470 the prior fluxes during the optimization. Emission ratios between NO2 and CO2 are also uncertain (those from the prior 471 are currently used). The atmospheric residence time of NO₂ is another major source of uncertainty. The inversion 472 reports total fossil CO₂ emissions calculated from NOx combustion emissions. However, in principle, the derivation 473 of CO₂ emissions from the NOx inversions should be restricted to fossil fuel CO₂ emissions based on the fossil fuel 474 CO₂/NOx ratio from the TNO, as there is a better-established relationship between CO₂ and NOx from combustion of 475 fossil fuels. Future inversions co-assimilating CO₂ data will make a clearer distinction in the processing of fossil-fuel 476 and other anthropogenic emissions. Finally, it's important to note that the inversion results are not fully independent 477 of the bottom-up methods, as the prior estimates are based on TNO gridded products. However, part of the lack of 478 departure from the prior can also be attributed to the general consistency between the prior and the observations, which 479 raise optimistic perspectives for the co-assimilation of co-emitted species with the data from future CO₂ networks 480 dedicated to anthropogenic emissions.

- 481
- 482 483

484 3.3. CO₂ land fluxes

485 This section updates the benchmark data collection of CO₂ emissions and removals from the LULUCF sector 486 in EU27+UK previously published in Petrescu et al. (2020) and Petrescu et al. (2021b), expanding on the scope of 487 those works by adding additional datasets and years. The countries analyzed in this study use country-specific activity





data and emissions factors for the most important land use categories and pools (EU NIR 2022, UK NIR 2022).
However, several gaps still exist, mainly in non-forest lands and non-biomass pools (e.g., EU NIR, 2022). In addition,
since NGHGIs largely rely on periodic forest inventories (carried out every five to ten years) for the most important
land use (Forest land), the net CO₂ LULUCF flux often does not capture the most recent changes, nor the full
interannual variability.

493 While the net LULUCF CO₂ flux was relatively stable from 1990 to 2016, staying mostly between -300 to -494 350 Mt CO₂/yr, in the past three years the sink has weakened to around -250 Mt CO₂/yr in 2020 (black dotted line in 495 Fig. B2, Appendix B; Abad-Viñas, pers. comm, 2022). This weakening occurred mostly in Forest land, due to a 496 combination of increased natural disturbances, forest aging and increased wood demand (Nabuurs et al., 2013; EU 497 NIR, 2022). Natural disturbances, including fires (especially in the southern Mediterranean), windthrows, droughts 498 and insect infestations (especially in central and northern European countries), have increased in recent years (e.g., 499 Seidl et al., 2014) which explains most of the interannual variability of NGHGIs. Forest aging affects the net sink 500 both through the forest growth (net increment) - which tends to level off or decline after a certain age - and the harvest, 501 because a greater area of forest reaches forest maturity (Grassi et al., 2018b). Although the exact increase in total 502 harvest in Europe in recent years is still subject to debate (Ceccherini et al., 2020; Palahi et al., 2021), demand 503 for fuelwood at least has increased (Camia et al., 2021).

504 Carbon uptake as seen by the atmosphere may occur on either managed or unmanaged land, and results from 505 processes such as photosynthesis, respiration, and disturbances (e.g., fire, pests, harvest). As discussed by Petrescu et 506 al. (2020), the fluxes reported in NGHGIs relate to emissions and removals from direct LULUCF activities (clearing 507 of vegetation for agricultural purposes, regrowth after agricultural abandonment, wood harvesting and recovery after 508 harvest and management) but also indirect CO₂ fluxes due to processes such as responses to environmental drivers on 509 managed land. Additional CO2 fluxes occur on unmanaged land, but these fluxes are very small in Europe. According 510 to Table 4.1 in the EU27 and UK NGHGIs (2021) CRF, almost all land (~95 %) in the EU27+UK is considered managed. France and Greece report some unmanaged forest areas (1.1 % and 16.6 %, respectively). Hungary and 511 512 Malta report unmanaged Grassland areas of 33 % and 100 %, respectively, and Nordic and Baltic countries plus 513 Ireland, Slovakia and Romania report sometimes quite large (up to 100 %) unmanaged wetland areas.

514 The indirect CO₂ fluxes on managed and unmanaged land due to changing climate, increasing atmospheric 515 carbon dioxide concentrations, and nitrogen deposition, are part of the (natural) land sink in the definition used in 516 IPCC Assessment Reports and the Global Carbon Project's annual global carbon budget (Friedlingstein et al., 2022), 517 while the direct LULUCF fluxes are termed "net land-use change flux", as discussed by Grassi et al. (2018a, 2021, 2022a), Petrescu et al. (2020, 2021b) and Pongratz et al. (2021). Results should thus be interpreted with caution due 518 519 to these definitional differences, but as most of the land in Europe is managed and the indirect effects are small, the 520 definitional differences should be modest compared to other sources of uncertainty (Petrescu et al., 2020). Other 521 relatively recent studies have already analyzed the European land carbon budget using GHG budgets from fluxes, inventories and inversions (Luyssaert et al., 2012) and from forest inventories (Pilli et al., 2017; Nabuurs et al., 2018). 522 523





524 3.3.2. LULUCF CO₂ fluxes from NGHGI and decadal changes

525

526 Figure 3 shows the decadal change in CO₂ LULUCF flux from the UNFCCC NGHGI (2019) (upper plot) 527 compared with the UNFCCC NGHGI (2021) (bottom plot). The contribution of each category ("remaining" and 528 "conversion") to the overall reduction of CO₂ emissions in percentages between the three mean periods (gray columns) 529 are the mean values over 1990-1999, 2000-2009 and 2010-(2017) 2019. The "+" and the "-" signs represent a source and a sink to the atmosphere. LUC(-) represents the land use conversion changes that increase the strength of the 530 531 LULUCF sink between two averages (i.e., values become more negative); LUC(+) represents the land use conversion changes that decrease the strength of the overall LULUCF sink. Note that the categories inside LUC(-) may be sources 532 533 or may be sinks, but between the two average periods, they become more negative. The HWP pool can constitute 534 either a source or a sink depending on the balance between the timber input to the pool (contributes to a sink) and the 535 loss of carbon as products reach their end-of-life (source). The absolute contributions of each category to the total LULUCF fluxes for 1990-2019 are given in Fig. B2 for context. 536

From the 1990–1999 mean to the 2000–2009 mean, the CO₂ LULUCF flux changed from -87.98 to -96.98 Tg C in the 2021 NGHGI (i.e., strengthened by 10.0 %), compared with -10.7 % for the 2019 NGHGI (note that Petrescu et al. (2021b) reported -9.6 %, which is the change relative to the 2000-2009 mean instead of the 1990-1999 mean that we adopt here due to common usage). This indicates a slight decrease in the reported European land sink compared to the previous estimates due to revised historical estimates. A 3.8 % growth in emissions from FL-FL and LUC(+) (Wetlands, Settlements and Other land conversions) weakened the overall sink¹⁰, while the sink related to all other categories grew by 15 % to strengthen the overall sink¹¹.

From the 2000-2009 mean to the 2010–2019 mean, the CO₂ LULUCF flux changed by +3.7 % (i.e., weakened sink), compared with +3.4 % reported by Petrescu et al., (2021b) which denotes a slight weakening of the European land sink compared to the previous estimate. Note the difference in time period (2010-2019 here, but 2010-2017 previously). A 9.6 % growth in emissions from FL-FL, HWP and LUC(+) (Forest land, Wetlands, and Settlements conversions) weakened the overall sink¹², while the sink related to all other categories changed by -5.9 % and strengthened the overall sink¹³.

550

¹⁰Positive percentages represent sources

¹¹Negative percentages represent sinks.

¹²Positive percentages represent sources

¹³Negative percentages represent sinks.







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552 Figure 3: The contribution of changes (%) in CO2 land fluxes from various LULUCF categories to the overall change 553 in decadal mean for the EU27+UK as reported by Member States to the UNFCCC. The top plot shows the previous 554 NGHGI data from Petrescu et al. (2021b) and the bottom plot illustrates data from UNFCCC NGHGI (2021). Changes 555 in land categories converted to other land are grouped to show net gains and net losses in the same column, with the 556 bar color dictating which category each emission belongs to; note that the composition of the "LUC(+)" and "LUC(-)" bars can change between time periods. Not shown are emissions from "Wetlands remaining wetlands", 557 558 "Settlements remaining settlements", and "Other land remaining other land" as none of the BU models used 559 distinguish these categories. The fluxes follow the atmospheric convention, where negative values represent a sink 560 while positive values represent a source. The color bars are shaded to guide the eye in the direction of the change 561 (white-to-color). 562

563 Similar to Petrescu et al. (2021b), changes of HWP emissions remain by far the major contributor to changes in 564 the LULUCF sink strength, but the direction of their contribution is opposite across the two periods: from 565 strengthening the sink during 1990–1999 to 2000–2009 to reducing the sink in 2010-2019. However, the balance





566 between HWP and FL-FL is quite sensitive to the periods selected: for the difference between 1993-2001 and 2002-567 2010, FL-FL contributes more (+7.3 %) than HWP (-5.2 %). EU-27+UK Member States have all implemented the 568 IPCC Approach B (i.e., production approach) for the HWP pool (EU NIR, 2021), which "inventories carbon in wood products from domestically harvested wood only and does not provide a complete inventory of wood carbon in 569 national stocks" (Volume 4, Chapter 12, IPCC, 2006). Figure 3 suggests that carbon emissions from HWP "end-of-570 571 life" became greater than the amount of carbon entering HWP from domestic harvest in recent decades. If the flux of 572 carbon into the HWP (a portion of domestic harvest) decreases, there will be a lag effect where outputs (due to wood 573 product end-of-life) may dominate, leading to a source from HWP. This is confirmed by a more detailed analysis of 574 the reported gains and losses for the bloc (see Figure A2 in the Appendix), which shows a drop in harvested wood 575 product gains around 2008 followed by a slower recovery compared to the pre-2008 trend. Gross losses from the 576 HWP pool, on the other hand, have been increasing as HWP produced pre-2008 reach their end-of-life, leading to a 577 weakened sink from 2009 onwards compared to during the mid-2000s.

578

579 3.3.3. Estimates of CO₂ land fluxes from bottom-up approaches

580 In this section we present annual total net CO₂ land emissions between 1990-2020 i.e., induced by both 581 LULUCF and natural processes (e.g. environmental changes) from class-specific models as well as from models that 582 simulate multiple land cover/land use classes. The definitions of the classes may differ from the IPCC definitions of 583 LULUCF (e.g., FL, CL, GL) where, according to IPCC 2006 guidelines, to become accountable in the NGHGI under 584 "remaining" categories, a land-use type must be in that class for at least N years (where N is the length of the transition 585 period; 20 years by default). In an effort to create the most accurate comparison as possible in terms of categories and 586 processes included, total Forest land (FL) has been divided up into Forest land remaining forest land (FL-FL) and 587 Land converted to forest land (X-FL), while only total Grassland (GL) and Cropland (CL) are reported. This is largely 588 due to the non-forest sector models explored here only considering net land use change, which prevents separating out 589 the "converted" component.

590

591 Forest land

592 Fluxes from Forest land which remain in this class (FL-FL) are shown in Fig. 4. These fluxes were simulated 593 with ecosystem models (CBM and EFISCEN-Space, described in more detail in the Appendices) and countries' official inventory statistics reported to UNFCCC. The results show that the differences between models are systematic, 594 595 with CBM having slightly weaker sinks than EFISCEN-Space. CBM updated its historical data (1990-2015) and 596 presents new NBP estimates based on extrapolation of historical timeseries (see Appendix A2) for 2017-2020 597 (CBMsim). Both CBM and EFISCEN-Space use national forest inventory (NFI) data as the main source of input to describe the current structure and composition of European forests. NFIs are also the main source of input data for 598 599 most countries in the EU27 for NGHGIs (EU NIR, 2021), including data for carbon stock changes in various pools as well as the estimation of forest areas. Given that EFISCEN-Space does not cover all countries in the EU27+UK 600 601 (Austria, Bulgaria, Denmark, Hungary, Lithuania, Portugal and Slovenia are missing), the results were scaled by 602 1/0.74 to account for the fact that the available countries comprise around 74 % of the forest NBP for the EU27+UK,





according to previous EFISCEN results (Petrescu et al., 2021b). As noted above, EU regulations are driving Member
 States to report spatially explicit NGHGIs. Unlike the original EFISCEN, EFISCEN-Space is a spatially explicit
 model, in addition to being able to simulate a wider variety of stand structures, species mixtures and management
 options. Note that EFISCEN-Space reports only a single mean value for forest fluxes from 2005-2020; the annually
 varying value shown in Fig. 4 arises from scaling by annually varying forest areas.

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Figure 4: Net CO₂ land flux from Forest land remaining forest land (FL-FL) estimates for EU27+UK CO₂ from the
Petrescu et al. (2021b) synthesis paper (top) and a comparable graph using the updated data this year (bottom).
Means are given for 2006-2015 (top) and 2005-2019 (bottom) on the right side of both plots. CBM FL-FL historical
estimates include 25 EU and UK countries (excl. Cyprus and Malta) and include new estimates for 2017-2020 (red





crosses). The relative error on the UNFCCC value represents the UNFCCC NGHGI (2021) MS-reported uncertainty
with no gap-filling (EU NIR, 2021). The fluxes follow the atmospheric convention, where negative values represent a

sink while positive values represent a source. Notice that some timeseries have been removed and placed in Fig. 5 as

617 some datasets more accurately depict fluxes from total Forest land (FL).

618 The UNFCCC NGHGI uncertainty of CO₂ estimates for FL-FL across the EU27+UK, computed with the error propagation method (95 % confidence interval) (IPCC, 2006), ranges between 34 % - 55 % when analyzed at 619 620 the country level for all years, as it varies as a function of the component fluxes (EU NIR, 2019). Despite contrasting 621 methodologies and input data for emission calculation and uncertainties in each method (Appendix A), there is 622 reasonable agreement on the trend in FL-FL fluxes from CBMsim and the UNFCCC NGHGI (2021) (Fig. 4). The 623 magnitude of the values between EFISCEN-Space and the NGHGI (2021) also agree well, though as noted above the 624 EFISCEN-Space results only vary with the amount of forest area which makes the trend much flatter. Given that all 625 three methods (NGHGI, CBM, and EFISCEN-Space) are heavily based on national forest inventory data, the general 626 agreement between the three is not surprising.

Figure 5 presents CO₂ land estimates for total Forest land (both remain and convert classes, "FL"). For the total Forest land, the results were simulated with an ecosystem model (ORCHIDEE) and a global dataset (FAOSTAT) as it is not possible for these two approaches to separate out the "remain" and "convert" land use category. This obstacle arises due to the use of net land use/land cover information which does not include detailed information on the nature of the conversions. Consequently, Fig. 5 compares them to the total Forest land from the countries' official inventory statistics (UNFCCC NGHGI, 2021).

633 From 2001 and until 2010, the FAOSTAT reports an increasing sink over time, which weakens from 2011 634 until 2019 (Fig. 5). This is explained by a reporting inconsistency in the Romanian inventory which had not been 635 corrected at the time of this analysis. Therefore, Romanian estimates for Forestland and Net forest conversion have been removed for the whole 1990-2020 timeseries in Fig. 5. Starting in 2016, FAOSTAT estimates better match those 636 from the NGHGIs as FAOSTAT updated its estimates. FAOSTAT uses input data directly from country submissions 637 to the FAO Global Forest Resource Assessments (FRA14) (e.g., carbon stock change is calculated by FAO directly 638 639 from carbon stocks and area data submitted by countries). It is important to note that these data are not always identical 640 to those submitted to the UNFCCC (Tubiello et al., 2021).

¹⁴The Global Forest Resource Assessment (*FRA*) is the supplementary source of Forest land data disseminated in *FAOSTAT*, http://www.FAO.org/forestry/fra/en/







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Figure 5: Net CO₂ land flux from total Forest land estimates (FL) for EU27+UK CO₂ from the UNFCCC NGHGI
(2021) submissions, the FAOSTAT data-driven inventory, and the ORCHIDEE DGVM. The relative error on the
UNFCCC value represents the UNFCCC NGHGI (2021) MS-reported uncertainty with no gap-filling (EU NIR, 2021).
FAOSTAT data does not include Romanian inventory estimates. The means are calculated for the 1990–2019
overlapping period. The fluxes follow the atmospheric convention, where negative values represent a sink while
positive values represent a source.

648 ORCHIDEE was updated to include a dynamic nitrogen cycle coupled to the carbon cycle in this work. As 649 shown in Appendix A2, the coupled nitrogen cycle results in a stronger sink, even if identical forcing is used. 650 ORCHIDEE shows a high inter-annual variability in carbon fluxes for forests in Fig. 5 because it incorporates 651 meteorological data at sub-monthly timescales, while methods based on forest inventories are generally updated only 652 every few years (e.g., five years for FRA), which results in a more climatological perspective. ORCHIDEE results 653 indicate that climatic perturbations and extreme events (multi-month droughts, in particular) can have significant 654 impacts on the net carbon fluxes depending on their timing in relation to the growing season. This is in line with flux 655 tower measurements that show significant year to year variability (Ciais et al. 2005). This is also to some extent 656 supported by dendrometer data although such data varies greatly among sites and tree species which obscures a 657 significant net effect (Scharnweber et al., 2020). It should also be noted that dendrometer data measures carbon stored 658 in individual trees, while the NBP reported in figures in this paper include fluxes from litter and soil respiration. The 659 variability of the weather data affects the carbon dynamics of all components of the ecosystems (hence NBP), which, 660 for instance, impacts on C assimilation rates, length of the growing season, dynamics of respiration rates and allocation of the carbon in the plant (cf. Fig. 1 and 2 in Reichstein et al. (2013) and Bastos et al. (2020b)). 661

A few reasons for differences between estimates seen in Fig. 4 and 5 can be readily identified. For this study,
the ORCHIDEE model used the ESA-CCI LUH2v2 PFT distribution (a combination of the ESA-CCI land cover map





664 for 2015 with the historical land cover reconstruction from LUH2 (Lurton et al., 2020)), and assumes that the shrub 665 land cover classes are equivalent to forest. In terms of area, the original ESA-CCI product corresponding to the 666 EU27+UK shows shrub land equal to about 50 % of the tree area in 2015. A similar analysis using the FAOSTAT 667 domain Land Cover, which maps and disseminates the areas of MODIS and ESA-CCI land cover classes to the SEEA 668 land cover categories15, shows that shrub-covered areas are around 20 % of that of forested areas for the 669 EU27+UK. The impact of classifying shrubs as "forests" on the total carbon fluxes could therefore account for a 670 significant percentage of the differences between ORCHIDEE and other results in Fig. 5. In addition, CBM depends 671 strongly on input data and related uncertainty. Historical data are retrieved from both country and EU statistics and 672 usually refers to forest management units rather than individual inventory plots. Finally, trends in forest carbon 673 strongly result from management, which are not represented in this version of ORCHIDEE but are included in CBM 674 and EFISCEN-Space.

675

676

677 Cropland

Cropland (CL, represented in the UNFCCC NGHGI 2021 as UNFCCC category 4B) includes net CO₂ emissions and removals from soil organic carbon (SOC) under "remaining" and "conversion" categories. Figure 6 shows the annual fluxes belonging to the category CL from the NGHGI for the EU27+UK along with four other approaches: one bottom-up inventory (FAOSTAT), two sector-specific models (EPIC-IIASA, ECOSSE), and one DGVM (ORCHIDEE). Note that the FAOSTAT value only includes the carbon flux from organic soils drained for agriculture, while ECOSSE, EPIC-IIASA, and ORCHIDEE include biomass volatilized immediately upon harvest; biomass left on site to decay as litter; and soil organic carbon.

The previous synthesis of Petrescu et al. (2021b) (Fig. 6, top) compared models against results for GL-GL from the NGHGI. For the current work, we compare against the total Grassland values (GL). The reason for this is that FAOSTAT, ECOSSE, EPIC-IIASA, and ORCHIDEE all use land use/land cover maps generated by IPCC Approach 1, which only records the total amount of land in a category for each year; information on transitions between categories is unknown. Therefore, it is not possible to separate out "remain" and "convert" categories.

For the common period (1990-2019), ORCHIDEE simulates a mean sink of -26 Tg C yr⁻¹, while ECOSSE, EPIC-IIASA, and FAOSTAT all simulate mean sources of 21 Tg C yr⁻¹, 10 Tg C yr⁻¹ and 16 Tg C yr⁻¹, respectively. With the exception of ORCHIDEE, all models are in line with the NGHGI results (mean over the same period of 22 Tg C yr⁻¹). In Petrescu et al. (2021b) (Fig. 6, top) the NGHGI reported a very small but constant source over the whole period (mean of 5.6 ± 3.5 Tg C yr⁻¹) with almost no inter-annual variability by construction, while all three processbased models simulated a sink.

The sink in ORCHIDEE must arise from the soil, as no simulated biomass in croplands remains from year to year; carbon is assimilated into biomass growth during the growing season, after which the biomass dies, is partitioned between litter and harvest (50 % to each), and either decays or vaporizes, respectively. In other words, no woody or perennial crops are simulated. NGHGIs assume that all aboveground biomass of non-woody crops re-enters the

¹⁵ http://www.fao.org/faostat/en/#data/LC





atmosphere at harvest. Given more favorable growing conditions due to climatic changes and CO₂ fertilization, this
 leads to more carbon entering the soil in ORCHIDEE in recent decades, which is driving the calculated CL sink
 observed in the model.

703 In the NGHGI, the reported source for the EU27+UK is mostly attributed to emissions from cropland on organic soils¹⁶ in the northern part of Europe where CO_2 is emitted due to C oxidation from tillage activities and 704 drainage of peat. The fact that FAOSTAT values are similar to the UNFCCC values points to the primary role of 705 706 drained organic soils, as this is the only flux included for the FAOSTAT dataset in Fig. 6. Finland and Sweden are of 707 particular importance, as they together account for more than half of the total area of organic soil in Europe. Organic 708 soils are an important source of emissions when they are under management practices that disturb the organic matter 709 stored in the soil. In general, the NGHGI emissions from these soils are reported using country-specific values when 710 they represent an important source within the total budget of GHG emissions.

711 ORCHIDEE also shows a much larger year-to-year variation due to the response of vegetation and respiration 712 fluxes to sub-daily meteorology. EPIC-IIASA and ECOSSE both operate on daily timescales (ECOSSE was updated 713 to daily for this work, though the previous version was monthly). As both photosynthesis (e.g., Kumarathunge et al., 714 2019) and respiration (e.g., Yvon-Durocher et al., 2012) show non-linear dependence on temperature, the more 715 extreme temperatures experienced by plants in ORCHIDEE will lead to a higher variation in vegetation response given 716 the same photosynthetic model. High IAV can be seen clearly for drought impacts in ORCHIDEE where regions change from sources to sink in a single year (e.g., for 2003 and 2018 (Ciais et al., 2005; Bastos et al., 2020a)). The 717 other two ecosystem models follow ORCHIDEE's patterns but with smaller magnitudes. FAOSTAT and NGHGIs 718 719 are mostly insensitive to inter-annual variability as the estimations are mainly based on statistical data for 720 surfaces/activities and emission factors that do not vary with changing environmental conditions.

721 Both ECOSSE and EPIC show a striking improvement in agreement with the NGHGI between V2019 (Fig. 6, top) and the current work (Fig. 6, bottom). For ECOSSE, this is the result of improved data, in particular around 722 723 residue management. The aboveground biomass is divided into harvest (which is accounted as direct emissions) and 724 residues (biomass that is partly removed and partly left on the field). The external tool MIAMI serves as the central 725 model for the NPP and follows the allocation distribution of Neumann and Smith (2018). The removed residues are 726 set to 50 % as a compromise between the wide range of residue removal rates given by Scarlat et al. (2010). Residue 727 and yield biomass from MIAMI are provided as input into the ECOSSE simulations. Additionally, more realistic 728 fertilizer data (Mueller et al., 2012) were used. For EPIC, the shifts in net CO₂ fluxes in the current EPIC results stem 729 from the updated soil organic carbon and nitrogen module (Balkovič et al., 2020) and updates in meteorological 730 forcing. Firstly, the updated soil module resulted in higher heterotrophic respiration across many EU regions. Besides 731 attributing more carbon to the soil surface emissions, enhanced respiration leads to higher NPP and yields in regions 732 with low fertilization rates as more nitrogen is released from the SOM pool. Secondly, altered solar radiation and air

¹⁶The 2006 IPCC Guidelines largely follow the definition of Histosols by the Food and Agriculture Organization (FAO), but have omitted the thickness criterion from the FAO definition to allow for often historically determined, country-specific definitions of organic soils (see Annex 3A.5, Chapter 3, Volume 4 of IPCC (2006) and Chapter 1, Section 1.2 (Note 3) of IPCC (2014)).





- 733 temperature data affected the full range of carbon variables in EPIC, including NPP, harvested biomass, heterotrophic
- respiration, and leached carbon.
- 735



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737 Figure 6: Net CO₂ land flux from Cropland estimates for the EU27+UK from: previous data from Petrescu et al., 738 (2021b) showing only the "remaining" fluxes (CL-CL) (top plot), and data from the UNFCCC NGHGI (2021) 739 submissions and models showing net carbon fluxes for the total Cropland (CL), with their 1990-2019 mean given on 740 the right (bottom plot). CL net carbon fluxes are estimated with three ecosystem models: ORCHIDEE, ECOSSE and EPIC-IIASA, in addition to the FAOSTAT inventory. Note that the FAOSTAT value only includes the carbon flux from 741 742 organic soils drained for agriculture. The relative error on the UNFCCC value represents the UNFCCC NGHGI 743 (2021) MS-reported uncertainty with no gap-filling (EU NIR, 2021). The fluxes follow the atmospheric convention, 744 where negative values represent a sink while positive values represent a source.



745

746 Finally, differences in the results between the models and the NGHGIs may arise from definitions. The 747 cropland definition in the IPCC includes cropping systems and agroforestry systems where vegetation falls below the 748 threshold used for the definition of Forest land category, consistent with the selection of national definitions (IPCC 749 glossary). Given that every country is allowed to select their definition of Forest land, which therefore influences the 750 area of Cropland and the total emissions, it is beyond the scope of this study to summarize here the criteria for the 28 751 countries under consideration and compare those to the methods used in determining the land use/land cover data for 752 the other models. However, the interested reader is referred to Tables 6.10 (forests), 6.18 (croplands), and 6.22 753 (grassland) in the 2022 NIR of the European Union (EEA/PUBL/2022/023).

754

755 Grassland

756 Grassland (GL, UNFCCC category 4C) includes net CO2 emissions and removals from soil organic carbon (SOC) under "remaining" and "conversion" categories. The grassland definition in the IPCC includes rangelands and 757 758 pasture land that is not considered as Cropland, as well as systems with vegetation that fall below the threshold used 759 in the Forest land category (same explanation as for Cropland). This category also includes all grassland from wild 760 lands to recreational areas as well as agricultural and silvo-pastoral systems, subdivided into managed and unmanaged, 761 consistent with national definitions (Petrescu et al., 2021b). For similar reasons to those expressed in the section Cropland above, the current work (Fig. 7, bottom) compares modeled CO₂ flux against NGHGI results for total 762 763 Grassland (GL).

The NGHGIs of countries in the EU27+UK report emissions from managed pastures and grasslands, although the details of what is included varies between countries (Table 6.21, EU NIR, 2021). Grasslands can be managed through grazing or by cutting. If a grassland is used for grazing but retains the natural vegetation, it is called a "rangeland". If the area has been replanted with vegetation specifically for animal forage, it is commonly referred to as "pasture"¹⁷. Since almost all European grasslands are somehow modified by human activity and to a major extent have been created and maintained by agricultural activities, they can be defined as "semi-natural grasslands", even if their plant communities are natural (Silva et al., 2008).

The NGHGI reports a slightly positive net flux over 1990-2019, although with a much larger uncertainty than 771 772 for either Forest land or Cropland (4 ± 28 Tg C yr⁻¹). While increased uncertainty compared to forest emissions is 773 understandable given the emphasis on collecting accurate forestry statistics due to their economic importance, the 774 increased uncertainty in Grassland compared to Cropland is more puzzling. Three possible explanations include: 1) 775 absolute Grassland emissions/removals are lower than for Cropland, which may lead to higher relative uncertainty 776 given the nearness to zero; 2) MS with lower uncertainties may dominate Cropland, while MS with higher 777 uncertainties may dominate Grassland; 3) Extensive work has been carried out on national/regional factors 778 representing changes in Cropland management, while less has been done on Grassland. For (3), this also may apply 779 to other biomass pools, as eight countries report "country specific" instead of "default" parameters for living biomass

¹⁷ See, for example, https://www.epa.gov/agriculture/agricultural-pasture-rangeland-and-grazing





in Cropland versus Grassland (while only one country does the reverse; Table 6.6., EU NIR, 2021). Additional
analysis will be needed to elucidate this issue.





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Figure 7: Net CO₂ land flux from total Grassland (GL) estimates for EU27+UK from: previous data from Petrescu et
al. (2021b) (top plot), and the updated datasets considered here (bottom plot). The means shown on the right of each
plot are for 1990-2017 (top) and 1990-2018 (bottom). GL net carbon fluxes are estimated with the ORCHIDEE, EPICIIASA, and ECOSSE (not updated and therefore identical to Petrescu et al., 2021b) models in addition to FAOSTAT.
The relative error on the UNFCCC value represents the UNFCCC NGHGI (2021) MS-reported uncertainty with no
gap-filling (EU NIR, 2021). The fluxes follow the atmospheric convention, where negative values represent a sink
while positive values represent a source.



793				
794	In addition to the NGHGI, updated results are available for ORCHIDEE (using a coupled C-N cycle) and			
795	FAOSTAT. For the first time, EPIC-IIASA contributed estimates for Grassland fluxes using five different grassland			
796	types and simulating carbon export due to herbivores (see Appendix A2 for more details). Both of these models			
797	exhibit a strong sink in Grassland. For ORCHIDEE, this is likely due to the same reasons as the sink in croplands:			
798	more suitable growing conditions due to climate change, CO2 fertilization, and nitrogen deposition leading to increased			
799	inputs into the soil which are not lost during tillage due to the lack of explicit management in the version reported			
800	here. For EPIC-IIASA, this results from manure left on site and incorporated into the soil. A Tier 1 IPCC approach			
801	assumes no changes in either living or dead biomass pools on grasslands; only considers organic soils which have			
802	been drained for grazing; and only considers mineral soils which have undergone a change in management. This			
803	greatly reduces or eliminates mechanisms which promote sinks in ORCHIDEE and EPIC-IIASA. On the other hand			
804	FAOSTAT reports a slight source in Grasslands, in line with the NGHGI. This is because, as is the case for Cropland,			
805	FAOSTAT data only considers emissions from drained organic soils. As incorporation of manure in EPIC-IIASA			
806	changes grasslands from a net source to a net sink, consideration of CO2 from manure input in other inventories may			
807	have a similar effect.			
808				
809 810	3.3.4. Bottom-up CO ₂ estimates from all LULUCF categories			
811	This section analyzes CO2 emissions and sinks for the LULUCF sector, including NGHGI categories (from			
812	Fig. 3) and a suite of different bottom-up approaches. This comparison is challenging due to differences in terms of			
813	activities covered in the different estimates, as well as differences in terminology (see, for example, Petrescu et al.,			
814	2020, Fig. 12). To summarize:			
815	• FAOSTAT differs from NGHGIs for reasons recently summarized by Tubiello et al. (2021), Petrescu et al.			
816	(2021b), and Grassi et al, (2022a), including numerically different data provided by Member States to			
817	FAOSTAT and UNFCCC; different methods (FAOSTAT applies a Tier 1 approach globally, while Member			

8 Member 818 States reports to the UNFCCC vary from Tier 1 to Tier 3); differences between net and gross land use change 819 (FAOSTAT is based on net transitions, following Approach 1 as detailed by the 2006 IPCC guidelines 820 (Chapter 3 of Volume 4, Sect. 3.3.1)); and differences in biomass pools. For the latter, FAOSTAT only considers living biomass pools instead of the five IPCC pools¹⁸ reported to the UNFCCC. A preliminary 821 examination shows that changes in dead wood, litter, and mineral soil carbon stock are generally less than 822 823 0.1 t C/ha, which is relatively small compared to reported changes around 1.0 t C/ha in living biomass pools 824 (Tables 6.13, 6.14, 6.15, EU NIR, 2021). On the other hand, changes in organic soil carbon stock are 825 approximately the same magnitude as living biomass, which may lead to significant discrepancies between 826 the NGHGI and FAOSTAT for the EU27.

¹⁸ According to the IPCC 2006 guidelines the reporting is done for the five LULUCF carbon pools: above-ground biomass, belowground biomass, dead wood, litter, and soil organic matter





- 827 DGVMs (represented here by the TRENDY v10 ensemble, as well as the high-resolution ORCHIDEE and 828 CABLE-POP simulations) include the impact of CO₂ fertilization, climate change and land use change for 829 Forest land, Grassland and Cropland categories; they do not explicitly treat the Wetlands, Settlement and 830 Other land categories as in the NGHGIs. They account for the evolution of living biomass, dead biomass, and soil organic carbon for all categories while for NGHGIs reporting is not mandatory for all subcategories 831 832 depending on the method Tier employed (e.g., dead organic matter in a Tier 1 method is assumed to be constant). There is significant uncertainty associated with the DGVMs' fluxes both from i) the forcing data, 833 834 including datasets of land-use changes and the coverage of different land use change practices, ii) model 835 parameters, and iii) model structural uncertainty (i.e., processes not included) (Arneth et al., 2017). Similar to FAOSTAT, DGVMs typically deal with net land use change emissions at the spatial resolution of the 836 837 model simulations (e.g., 0.5° or 1° for the TRENDY ensemble and 0.125° for the ORCHIDEE and CABLE-838 POP simulations) instead of gross land use change as reported in NGHGIs. CABLE-POP is an exception to 839 most DGVMs and actually incorporates gross land use transitions (Haverd et al., 2018). The use of gross land use transitions may induce significant differences with coarse resolution model simulations (e.g., the 840 841 TRENDY ensemble). In addition, DGVMs often do not distinguish between managed and unmanaged land, 842 while NGHGIs report results only from managed land.
- 843 The bookkeeping models, BLUE and H&N, calculate net emissions from land use change including 844 immediate emissions following land conversion, legacy emissions from slash and soil carbon decomposition after land-use change, carbon uptake during regrowth of secondary forest after pasture and cropland 845 846 abandonment, and emissions from harvested wood products as they decay. While activities on the category 847 Land remaining land are generally not considered in bookkeeping models, one major exception is fluxes from 848 wood harvest, which are a primary source of emissions on managed forest land. In addition, bookkeeping 849 models do not account for fluxes arising from "indirect" anthropogenic influences such as CO2 fertilization 850 or climate change.
- 851

852 Given all these differences in terms of activities, the comparison in this section should be considered as a 853 rough overview that highlights both important aspects of the C cycle and questions that need to be addressed in the 854 future. Going towards a more specific comparison of only net land-use change (LUC) fluxes would require additional 855 considerations. In GCP's annual global carbon budget, net LUC term is estimated by global DGVMs as the difference 856 between a run with and a run without land-use change (i.e., the S3 and S2 simulations from TRENDY, respectively) 857 and by bookkeeping models (Friedlingstein et al., 2022). Such an estimate is given in Fig. 13 in Petrescu et al. (2020) 858 for Forest land. However, this approach does not fully resolve the differences mentioned above. In particular, 859 questions remain about net vs. gross land use change, managed vs. unmanaged land, and emissions from wood harvest. In addition, UNFCCC "convert" emissions (i.e., emissions resulting from land that has been converted from 860 861 one type to another) are reported within 20 years following conversion in the "convert" category (biomass losses are 862 typically reported in the year of conversion, while net changes in soil organic carbon during the entire conversion 863 period). FAOSTAT, DGVMs, and bookkeeping models usually only include "convert" fluxes from the year following





864 conversion, although bookkeeping models and DGVMs which deal with gross transitions may be able to include this

865 transition period more easily.





EU27+UK: Net bottom-up CO₂ land fluxes from land use, land use change, and forestry (LULUCF)

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867 Figure 8: Net CO₂ fluxes from total LULUCF activities in the EU27 + UK from previous data from Petrescu et al. (2021b) (top plot) and data from seven new or updated sources (bottom plot) including: UNFCCC NGHGI (2021), 868 869 BLUE (vVERIFY), BLUE (vGCP2021), H&N (GCP2021), DGVMs (TRENDY v10), FAOSTAT (2021), ORCHIDEE 870 and CABLE-POP with high-spatial-resolution (0.125°) meteorological forcing (both models are also part of the 871 TRENDY ensemble at 0.5°). The gray bars represent the individual model data for the DGVMs. The UNFCCC estimate 872 includes all classes (remain and convert), as well as HWP. The relative error of the UNFCCC values represent the





WNFCCC NGHGI (2021) Member States reported uncertainty computed with the error propagation method (95% confidence interval), gap-filled and provided for each year of the timeseries. Biomass burning emissions are included
in the C stock estimates. The FAOSTAT estimate includes both Forest land remaining forest land in addition to
incorporating afforestation and deforestation as conversion of Forest land to other land types. The means are
calculated for the 1990–2019 overlapping period. The fluxes follow the atmospheric convention, where negative
values represent a sink while positive values represent a source.

879 Figure 8 shows CO₂ fluxes from the NGHGI LULUCF sector compared to all other comparable bottom-up 880 (BU) estimates in this work: high-resolution S3 simulations for both ORCHIDEE and CABLE-POP; the median of 881 15 S3 simulations from the TRENDYv10 DGVM ensemble; three bookkeeping models; and FAOSTAT. As 882 mentioned above, taking the difference of the TRENDY S2 and S3 simulations provides an estimate of the net flux 883 from land use change, but inconsistencies are introduced either way, and therefore further research is needed in order 884 to establish which approach (S3-S2, or simply S3) leads to the most consistent comparison. For the overlapping period 1990-2019, the means of two out of the three bookkeeping models (BLUE vGCP (-61 Tg C yr⁻¹) and BLUE vVERIFY 885 (-43 Tg C yr⁻¹, using the Hilda+ land use forcing)) along with the mean of FAOSTAT (without Romania) (-93 Tg C 886 yr¹) fall within the 95 % confidence interval of the UNFCCC NGHGI estimate of -86 ± 33 Tg C yr¹. Only H&N rests 887 888 apart with a stronger sink (-142 Tg C yr⁻¹).

889 Bookkeeping models like BLUE and H&N do not include indirect effects on biomass growth due to factors 890 such as CO2 fertilization, nitrogen deposition, and climate change, while NGHGIs implicitly include these impacts on 891 managed land through updated statistics. Recent work by Grassi et al. (2022b) demonstrates that including the sink 892 associated with human-induced indirect effects (as estimated by the S2 simulations from the TRENDY DGVM 893 ensemble) into results by bookkeeping models can largely reconcile estimates of net global LULUCF fluxes between 894 the NGHGIs and bookkeeping models. At the level of the EU27+UK, the inclusion of this sink results in an overcompensation; the BMs estimate a net sink of -56.5 Tg C yr⁻¹ compared to the NGHGI estimate of -87.9 Tg C yr⁻¹ 895 896 ¹, while the BMs+DGVMs results in -112 Tg C yr⁻¹. However, all of these estimates fall inside the NGHGI uncertainty 897 range in Fig. 8. This suggests that indirect effects are small in the EU27+UK.

898 The UNFCCC LULUCF estimates contain CO2 emissions from all six land use categories and HWP, 899 including remaining categories and conversion to and from a category to another. The DGVMs show high interannual 900 variability, as demonstrated clearly by the high-resolution CABLE-POP simulation in Fig. 8. The mean values for 901 DGVMs across the overlapping period, on the other hand, agree fairly well with the NGHGI: -170 Tg C yr⁻¹, -84 Tg C yr⁻¹, and -81 (min -285, max 118) Tg C yr⁻¹ for ORCHIDEE, CABLE-POP, and TRENDY v10, respectively, 902 903 compared to the NGHGI mean of -86 ± 33 Tg C yr⁻¹. Note again that ORCHIDEE and CABLE-POP are also part of 904 the TRENDYv10 ensemble, but the simulations included in TRENDY used a coarser meteorological forcing than the 905 one used within the VERIFY project (around 0.125° resolution). CABLE-POP also used a higher resolution land use 906 land cover change (LULCC) dataset for the results submitted to VERIFY (0.25° as opposed to 1.0°). The increased 907 IAV from the high-resolution CABLE-POP compared to ORCHIDEE is suspected to have been introduced through 908 the construction of the LULCC dataset as described in Appendix A2. Gross fluxes are, by definition, larger than net





909 fluxes, and consequently a method which incorporates gross fluxes (like CABLE-POP) can be expected to undergo910 larger changes than a method incorporating net fluxes (like ORCHIDEE).

911 The differences between bookkeeping models and UNFCCC and FAOSTAT are discussed in detail 912 elsewhere, and focus on the inclusion of unmanaged land in bookkeeping models but not FAOSTAT and UNFCCC methodologies (Petrescu et al., 2020; Grassi et al., 2018a, 2021). ORCHIDEE, CABLE-POP and the TRENDY v10 913 914 ensemble means show much higher inter-annual variability due to the sensitivity of the model fluxes to highly variable 915 meteorological forcing at sub-daily time steps which allow for much more rapid responses to changing conditions, as 916 already discussed in the previous sections. The incorporation of variable climate data and the fact that DGVM models 917 simulate explicitly climate impacts on CO₂ fluxes, which inventories and bookkeeping models do not, explain these 918 differences. A comparison including sector-specific models (e.g., ECOSSE, EFISCEN-Space, EPIC-IIASA, CBM) 919 where multiple model results are harmonized and aggregated to produce a "total" LULUCF flux comparable to 920 DGVMs and bookkeeping models would be insightful; however, such a comparison requires extensive analysis which 921 is beyond the scope of the current work.

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23 3.3.5. Comparison of atmospheric inversions with NGHGI CO₂ estimates

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925 Figure 9 highlights the range of estimates from global and regional atmospheric inversions (GCP2021, 926 EUROCOM, CSR, LUMIA, and CIF-CHIMERE; see Table 2 and Appendix A2 for more details) against bottom-up 927 total annual EU27+UK CO₂ land emissions/removals from the UNFCCC NGHGI (2021). The top panel in the figure 928 shows the previous results from Petrescu et al. (2021b). In these inversions, all components of the carbon cycle that 929 contribute to the observed atmospheric CO₂ gradients between stations are implicitly included as the inversions incorporate observed atmospheric concentrations of CO2. This includes processes where carbon is uptaken by 930 931 vegetation in one area and emitted in a different area, i.e. emissions due to the respiration of laterally transported 932 carbon.

933 One significant change between this work and Petrescu et al. (2021b) is the removal of emissions and sinks 934 from inversion results due to lateral transport of carbon from crop trade, wood trade, and inland waters. Bottom-up 935 methods (including all the NGHGIs for European countries) do not consider emissions and removal of atmospheric 936 CO2 due to lateral transport of carbon, while observations assimilated into top-down inversions record all CO2 fluxes 937 without separating their components. We followed Eq. (1) of Deng et al. (2021) without prior masking for managed 938 land. Emissions from lateral transport of carbon ("lateral fluxes") were prepared generally following the approach 939 described by Ciais et al. (2021), where crop and wood product fluxes are derived from country-level trade statistics 940 compiled by the FAO. Inland water emissions and riverine export of terrestrial carbon use spatially explicit 941 climatological data and a statistical model combined with estimates of gas transfer velocities. A more complete 942 description is given in Appendix A2. This adjustment has been applied to all top-down fluxes reported here unless 943 indicated otherwise.

944 The C fluxes from inland waters (rivers and lakes) reported in Petrescu et al. (2021b), were replaced in this
945 study by maps of sinks/sources of rivers/lakes, wood and crops, accounting for a combined mean of -136 Tg C yr⁻¹
946 (over the 2010-2018 common period of the inversions). For comparing bottom-up methods (including the NGHGI) to





947 TD estimates in the EU27+UK, it is always necessary to remove the traded wood and crop harvest (see Deng et al. 948 (2021) for additional explanations). For the NGHGI, this arises due to how harvested wood products are considered. 949 HWPs can be reported to the UNFCCC by multiple approaches, three of which are outlined in Chapter 12 of Volume 950 4 of the 2006 IPCC Guidelines. One of these methods (the Atmospheric Flow Approach) would allow for a direct 951 comparison with the inversions as wood product emissions are accounted for in countries in which they are in use and 952 in landfills. However, all countries in the EU27 adopt the Production Approach (2022 NIR of the European Union 953 (EEA/PUBL/2022/023)) in which emissions are considered due to domestic harvest regardless of where the wood is 954 transformed or used. Inversions, on the other hand, see the HWPs where they transform into CO₂, either through 955 decomposition or incineration. It should be noted that DGVMs also typically implement the Production Approach on 956 a pixel level (i.e., harvested wood decomposes in the pixel where it is produced). As pixels reported for the high-957 resolution simulations here are around 10 km wide, this implicitly assumes that HWP never travel more than 10 km 958 from the harvest site (this becomes 50 km in coaster resolution simulations like TRENDY). Therefore, removing 959 emissions from lateral carbon transport makes inversions more comparable not only to NFGHGIs but also to DGVMs.

960 Flux estimates from inversion methods for CO₂ land show much more variability than the NGHGI (Fig. 9). 961 The mean of the EUROCOM ensemble of European inversions shows good agreement with UNFCCC NGHGI data, 962 but with a huge spread of annual model results that extends from significant sources into large sinks. This large spread 963 can be linked to uncertainty in atmospheric transport modeling, inversion methods and assumptions, and to limitations 964 of the observation system. Furthermore, the EUROCOM inversions were designed for the European geographical 965 domain (which is larger than the EU27+UK) and are still being developed in particular to better constrain the 966 latitudinal and longitudinal boundary conditions.

967The annual mean (overlapping period 2010-2018) of the EUROCOM v2021 inversions (-80 [-175,-4] Tg C968 yr^{-1}) is the closest inversion estimate to the timeseries mean of the NGHGI estimates (-88 ± 31 Tg C yr⁻¹), where the969error bars for the inversion indicated the [0th,100th] percentiles due to the small size of the ensembles. The mean of970the global GCP2021 inversions (-50 [-320,+122] Tg C yr⁻¹) and regional inversions, CSR (-46 [-126,+47] Tg C yr⁻¹)971and LUMIA (-65 [-97,-27] Tg C yr⁻¹) show a lower absolute value, but report larger interannual variability (min/max).972The new CIF-CIMERE product has a mean of -99 Tg C yr⁻¹, showing more negative fluxes since 2010, which is not973seen in other models and is still under investigation.

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Figure 9: Comparison of inventories and atmospheric inversions for the total EU27+UK biogenic CO₂ fluxes from
Petrescu et al. (2021b) (top plot) and updated data from current study (bottom plot). Top-down inversion results are:
the global GCB2021 ensemble, the regional EUROCOM ensemble, the regional CarboScopeReg model with multiple
variants, the regional LUMIA model with multiple variants, and CIF-CHIMERE. The relative error in the UNFCCC





981 values represents the UNFCCC NGHGI (2021) Member states reported uncertainty computed with the error propagation method (95% confidence interval) gap-filled and provided for every year of the timeseries. The timeseries 982 983 mean overlapping period is 2010-2018. The colored area represents the min/max of model ensemble estimates. The 984 same emissions due to lateral fluxes of carbon through rivers, crop trade, and wood trade are removed from the top-985 down estimates in both the top and bottom graphs for consistency. The fluxes follow the atmospheric convention, 986 where negative values represent a sink while positive values represent a source. Note that Petrescu et al. (2021b) 987 presented the top plot including a suite of bottom-up models, which have been removed here for clarity as they have 988 already been presented in Fig. 8.

989

The comparison of past and current versions of the inversions shows changes in specific models. A reduction
in the spread of the estimates is noted over the two past versions of CSR, resulting in a small source in the most recent
estimates. The CSRv2021 (bottom-plot) predicts in 2018 (last common year of both versions) a small source of 19 [64, +100] Tg C yr⁻¹ compared to the previous CSRv2019 which simulated a very strong sink of -253 [-280, -194] Tg
C yr⁻¹. This smaller source appears more in line with more positive fluxes expected in years of extreme drought (e.g.,
2018 in Northern Europe, although this did not impact the whole EU27+UK (Toreti et al., 2019)).

996As can be seen in Fig. 9, there is also improved agreement between the EUROCOM ensemble and the997NGHGI, including a greatly reduced IAV compared to the previous version. The small EUROCOM ensemble mean998sink for the 2009-2015 period of -1.9 [-335,+322] Tg C yr⁻¹ (top panel) strengthened to -93 [-187,-15] Tg C yr⁻¹ in the999v2021 version (bottom panel). The UNFCCC total LULUCF mean is -92 ± 33 Tg C yr⁻¹ for the same time period. The1000IAV of EUROCOM was dramatically reduced by removing the FLEXINVERT model from the v2021 ensemble as a1001clear outlier of annual means due to a slightly shifted seasonal cycle (Appendix A2).

The new GCP2021 inversions show a clear trend towards decreasing the CO₂ sink strength of the land surface after 2017, contrary to the NGHGI estimates which are relatively stable (Fig. 9, bottom). The large variability and high sink observed in the upper plot of Fig. 9 shifted to a source in 2019 (21 [-185, +226] Tg C yr⁻¹) due to the extreme climatic response of the TD models to the drought year, which can also be observed in the BU simulations (e.g., TRENDY v10, ORCHIDEE, and CABLE-POP in Fig. 8). Out of the GCP2021 models, CAMS was the model responsible for the lower sinks (data not shown), which may be due partly to changes in the stations assimilated.

1008 Table B2 summarizes the processes included in the CO₂ land models presented in this work, as these 1009 processes are seen for the moment as the main cause of discrepancies between estimates shown in all the previous 1010 figures. According to Table B2, no bottom-up model or dataset used here contains all of the 13 LULUCF categories 1011 reported in the NGHGIs. A simple analysis of the mean 1990-2020 LULUCF fluxes from the EU27+UK NGHGI 1012 (Table A3 in Appendix A2) shows that six categories account for almost 90 % of the gross flux: Forest land remaining 1013 forest land (56 %), Land converted to cropland (7 %), Land converted to forest land (7 %), Grassland remaining grassland (6 %), Harvested wood products (6 %), and Land converted to settlements (6 %). DGVMs currently include 1014 more of these categories than other methods. As shown in Fig. 8, the mean 1990-2019 value of the mean of the 15 1015 1016 TRENDY DGVM simulations is -81.9 Tg C yr-1 (with a a range of [-285,118] Tg C yr⁻¹), while those of the 1017 ORCHIDEE and CABLE-POP simulations using the high-resolution forcing provided in the VERIFY project are -





1018 171 Tg C yr⁻¹ and -84.8 Tg C yr⁻¹, respectively. The means agree quite well for TRENDYv10 and CABLE-POP, but 1019 the spread of all the DGVMs is quite large. In addition, the number of categories included may not be a good proxy 1020 for quality of comparison. While an ideal model would include all categories in the NGHGI, it must also represent 1021 these categories well. Figures 4-7 suggest that sector-specific models currently show better agreement with the 1022 NGHGI than DGVMs, although a more detailed analysis including the entire suite of TRENDY models would be 1023 insightful. Note that these categories are used as input to top-down approaches, and therefore cannot be disaggregated 1024 into results after the simulation.

1025 3.3.6. Uncertainties in top-down and bottom-up estimates

1026 Uncertainties are essential for complete comparisons between models and approaches. This section 1027 summarizes the main sources of uncertainty estimates interwoven throughout the above text. We also provide a 1028 comparison of available uncertainties between the previous synthesis (V2019) and the current synthesis (V2021) for 1029 both bottom-up and top-down methods. Finally, we give an overview of two important advances in uncertainty 1030 estimation included in this work (one for the NGHGI, and one for top-down approaches), referring the interested 1031 reader to the Appendix for more information.

Several sources of uncertainty arise from the synthesis of bottom-up (BU) inventories and models of carbon fluxes, which can be summarized as: (a) differences due to input data and structural/parametric uncertainty of models (Houghton et al., 2012) and (b) differences in definitions (Pongratz et al., 2014; Grassi et al., 2018b, 2021; Petrescu et al., 2020, 2021b). Posterior uncertainties in top-down (TD) estimates mostly come from: 1) errors in the modeled atmospheric transport; 2) aggregation errors, i.e., errors arising from the way the flux variables are discretized in space and time and error correlations in time; 3) errors in the background mole fractions; and 4) incomplete information from the observations and hence the dependence on the prior fluxes.

1039 Figure 10 summarizes the quantifiable uncertainties in this work, compared to previous results from Petrescu 1040 et al. (2021b). With the exception of the NGHGI, all the other uncertainties are calculated from ensembles of 1041 simulations using either: 1) multiple models of the same general type, either using model-specific inputs or attempting 1042 to harmonize inputs as much as possible (e.g., TRENDY), or 2) multiple simulations with the same model, varying 1043 input parameters and/or forcing data (e.g., CarboScopeRegional, LUMIA). As a complete characterization of model 1044 uncertainty involves exploring the full parameter, input data, and model structure space, none of the uncertainties 1045 reported here can be considered "complete", but they represent best estimates given realistic constraints of resources 1046 and knowledge. The uncertainties represent the mean of overlapping periods for the previous V2019 (overlapping 1047 period: 2006-2015) versus the current V2021 (2010-2018). In general, the differences in mean behaviors between the 1048 two versions falls within uncertainty estimates. Note, however, that this graph can hide certain behaviors. For example, the similarity in the means for ORCHIDEE-VERIFY for both periods (-128.5 and -131.0 Tg C yr⁻¹ for V2019 1049 1050 and V2021, respectively) is likely a coincidence, given the wide fluctuation of annual values and the differences in 1051 the multi-decennial means seen in Fig. 8.

1052Figure 10 shows notable reductions in the spread of two ensembles: EUROCOM and CSR. Both of these1053are regional ensembles. In addition, the CSR results show a weaker sink in the current V2021 version compared to1054the previous V2019 version. As noted in Appendix A2, the change for CSR is explained by the inclusion of a corrected





- 1055 observation dataset for an isolated station in southeastern Europe which heavily influenced the regional results. The
- 1056 reduction in the spread of the EUROCOM ensemble results from the exclusion of a single member which produces
- annual flux results that are clear outliers compared to the remaining three members. More details of this analysis can
- 1058 be found in Appendix A2. The remaining ensembles retain similar model spread compared to the previous versions.



Mean of overlapping timeseries - V2019 (2006-2015) and V2021 (2010-2018) EU27+UK : net land CO_2 fluxes

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1060 Figure 10: Mean annual values of overlapping time periods (2006-2015) from Petrescu et al. (2021b) (transparent 1061 boxes and light gray lines) and new means for the 2010-2018 period from the current study (Fig. 8 and 9, Sect. 3.3.4 and 3.3.5). The hashed boxes and colored boxes depict the "old" and "new" values for ensembles of multiple models, 1062 1063 with the top and bottom of the boxes corresponding to minimum and maximum mean values of the overlapping period. 1064 For non-ensemble models (e.g., CIF-CHIMERE, FAOSTAT) the mean of the old and new overlapping periods are 1065 given by gray dotted and black dashed lines, respectively. The NGHGI UNFCCC uncertainty is calculated for 1066 submission year 2021 as the relative error of the NGHGI value, computed with the 95 % confidence interval method gap-filled and provided for every year of the timeseries. Inversions for both V2019 and V2021 have been corrected 1067 1068 for emissions of CO₂ from lateral transport of carbon using identical datasets to enable a fair comparison. The fluxes 1069 follow the atmospheric convention, where negative values represent a sink while positive values represent a source. 1070

1071 Three advances in uncertainty estimation were made in this study, involving all three classes of models: 1072 NGHGI, bottom-up, and top-down. In Petrescu et al. (2021b), percentage uncertainties for the NGHGI (2019) 1073 LULUCF sector and land use categories were taken from reported uncertainties of the EU Member States and UK that 1074 are used for compiling the National Inventory Reports (NIR) of the EU27+UK bloc, as well as the aggregate 1075 uncertainties for the block reported in the EU NIR. Uncertainty estimates were only given for a single year and were





1076 also partially incomplete due to missing uncertainty estimates for some sectors/subsectors of some countries. For the 1077 current work, we use values compiled by the EU inventory team involving a recently developed procedure to 1078 harmonize and gap-fill uncertainties reported by the Member States at the sector level (see EU NIR, 2021). Error 1079 correlations are accounted for, in addition to year-to-year variations in sub-sectoral contributions to the overall 1080 uncertainty. Extensive details are found in Appendix A1, and permit estimates of uncertainty on an annual basis, as 1081 opposed to the single value used in the previous synthesis. Note, however, that this procedure was not applied to sub-1082 sectoral categories (FL, CL, and GL), for which values were taken directly from EU NIR (2021) and applied across 1083 the whole timeseries. Synthesis plots created for individual countries and reported on the VERIFY website (VERIFY 1084 Synthesis Plots, 2022) take percentages directly from the respective country's NIR.

1085 The second advance relates to the impact of forcing data on bottom up models, in particular DGVMs. Figure 1086 A3 (Appendix A) shows how the ORCHIDEE model responds to both changes in meteorological forcing (for 1087 ORCHIDEE) and nitrogen forcing (for ORCHIDEE-N) over the past several decades. The impact of both is relatively 1088 small compared to interannual variability. This is likely due to at least two reasons. The first reason is that 1089 meteorological forcing used in this work has been re-aligned to the CRU observational dataset at 0.5 degrees and 1090 monthly resolution, thus removing large-scale and long-term differences between the original meteorological datasets. 1091 In addition, extensive spin-up and transient simulations are run for ORCHIDEE before reaching the point at which 1092 the forcing changes (1981 for the meteorological forcing, and 1995 for the nitrogen forcing). Such lengthy simulations 1093 enable woody biomass and soil carbon pools to develop a significant amount of inertia in response to additional 1094 changes. Greater differences may be seen for models where modified forcing data covers the entire length of the pre-1095 production simulation steps.

1096 The final advance relates to uncertainty characterization in the regional inversion model CSR following the 1097 methodology of Chevallier et al. (2007). Spatially explicit estimates of the uncertainty reduction achieved from the 1098 flux optimization were prepared through a Monte Carlo approach using an ensemble of 40 members. The uncertainty 1099 reduction is then calculated based on the ratio of the prior errors and the posterior spread of the ensemble members, 1100 using a formula such that 0 indicates no reduction and 1 indicates a complete elimination of uncertainty. A preliminary 1101 analysis showed that a considerable reduction may be achieved through the inclusion of more observation stations, 1102 although additional work is needed. For the moment, these maps only reflect random uncertainties, and systematic 1103 uncertainties remain poorly characterized. More information can be found in Appendix A2.

1104 Figure 11 presents an idea of the spatial uncertainties associated with these datasets. Total CO₂ land fluxes 1105 from EU27+UK and five main regions in Europe are presented, divided into top-down (top panel) and bottom-up 1106 (bottom panel) approaches for clarity. The regions (North, West, Central, East and South) consist of Annex I Parties 1107 to UNFCCC both inside and outside of the EU27+UK bloc, and are listed in Table A1. Figure 11 shows the total CO2 1108 land fluxes from the NGHGIs for base year 1990, as well as five-year mean values for the 2011-2015 and 2015-2019 1109 periods. The five-year periods are used as an exercise for what could be achieved in the first GST and also because 1110 they provided the most overlap with the datasets reported here. As the BU models in VERIFY include and simulate 1111 CO2 fluxes for at most three out of the six classes reported to the UNFCCC (FL, CL and GL), for comparison and 1112 consistency purposes both UNFCCC total LULUCF (including all six classes and HWP), as well as the UNFCCC





FL+CL+GL estimates are shown. Figure 11 presents CO₂ fluxes that include both direct and indirect LULUCF effects
on managed land. The total UNFCCC estimates include the total LULUCF emissions and sinks (by the UNFCCC
definition) belonging to all six IPCC land classes and the HWP class (see Sect. 2.3 and Appendix B for more details).
The NGHGI estimates are plotted and compared against fluxes simulated with statistical global and regional datasets:
bookkeeping models, biosphere and sector-specific models, and inversion model ensembles. The error bar represents
the variability in model estimates as the min and max values in the ensemble.









1121 Figure 11: Five-year means (2011–2015 and 2015-2019 as hashed and colored bars, respectively) of total CO₂ land

- 1122 flux estimates (in Tg C) for EU27+UK and five European regions (North, West, Central, South and East) for top-
- 1123 down (top) and bottom-up (bottom) methods compared to inventories. Eastern European region does not include
- 1124European Russia. Northern Europe includes Norway. Central Europe includes Switzerland. The UNFCCC
- 1125 uncertainty for the Republic of Moldova was not available. The data comes from: UNFCCC NGHGI (2021) total





1126 LULUCF submissions (dark green) which are plotted with respective base year 1990 (black star) estimates, the 1127 UNFCCC NGHGI (2021) FL+CL+GL estimates (light green), sector-specific BU models for FL, CL and GL (CBM, 1128 EPIC-IIASA, ECOSSE), ecosystem models (ORCHIDEE, TRENDY v10 DGVMs, CABLE-POP), global dataset 1129 FAOSTAT, bookkeeping models (BLUE (vGCP, and vVERIFY) and H&N), total CO₂ flux from TD inversion 1130 ensembles (GCP2021, EUROCOM) and three regional European inversions (CarboScopeReg (CSR), LUMIA and 1131 CIF-CHIMERE). ECOSSE GL data was not updated beyond 2018. Lateral CO2 fluxes (rivers/lakes, wood and 1132 crops sinks/sources) are represented separately (orange) and are removed from the top-down estimates as explained 1133 in the text. The fluxes follow the atmospheric convention, where negative values represent a sink while positive 1134 values represent a source. 1135 In general across the regions, BU (observation-based and process-based models) agree well with the 1136 UNFCCC-reported total LULUCF sources and sinks, except for the CABLE-POP DGVM which simulates a source 1137 for Central and Western Europe. As can be seen from the figure, however, this is not unexpected; the ensemble of 1138 TRENDY DGVMs shows a very large spread, and as such some DGVMs will undoubtedly display more extreme 1139 behavior. There remain however large disagreements between all estimates for Eastern Europe. This could be related 1140 to reduced data coverage for this region, in particular for the top-down approaches which depend on atmospheric measurement stations. In Northern Europe, some inversions agree with the NGHGIs on the magnitude of the sink 1141 1142 (mean of 2015-2019 of -65 Tg C yr⁻¹), while in Central Europe there is a large variance between the models. The 1143 differences are explained by updates and methodological changes detailed in Sect. 3.3.2 (sector specific process-based 1144 models and NGHGI), 3.3.3 (DGVMs, bookkeeping models and NGHGI) and 3.3.4 (all BU, TD and NGHGI). Finally,

the TD estimates are better in line with the NGHGI and the BU estimates after the removal of emissions due to lateral fluxes of carbon (discussed in Sect. 3.3.4). However, large variations still remain in the range of min/max of model ensembles represented in the figure by the error bars. For some models with high inter-annual variability (e.g., CIF-CHIMERE and CABLE-POP), the five-mean changes drastically between the two time periods but this may not represent a significant trend.

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1151 4. Data availability

Annual timeseries for the EU27+UK used in creation of the figures in this work for V2019 and V2021 are available for public download at <u>https://doi.org/10.5281/zenodo.7365863</u> (McGrath et al., 2022). This excludes CO₂ fossil data for the IEA, which is subject to license restrictions. The data are reachable with one click (without the need for entering login and password), and downloadable with a second click, consistent with the two click access principle for data published in ESSD (Carlson and Oda, 2018). The data and the DOI number are subject to future updates and only refers to this version of the paper. In addition, figures and annual timeseries for EU27+UK as well as other countries and regions are available from VERIFY Synthesis Plots (2022).

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1160 5. Summary and concluding remarks

This work represents an update of the Petrescu et al. (2021b) European CO₂ synthesis paper presenting and
 investigating differences between the UNFCCC NGHGI, BU data-based inventories, both coarse and high resolution

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1163 process-based BU models, and TD approaches represented by both global and regional inversions. Datasets used in 1164 the previous work have been updated by extending the temporal coverage and updating the models and data behind 1165 the calculations. In addition, several new models to expand the number of independent approaches compared have 1166 been added. Additional efforts have been made to improve uncertainty characterization in two approaches, along with 1167 a first attempt to present as many datasets as possible in a clear single figure to draw overarching conclusions.

1168 CO₂ fossil emissions dominate the anthropogenic CO₂ flux in the EU27+UK, regardless of the approach 1169 employed and irrespective of uncertainties. Fossil CO2 emissions are more straightforward to estimate than ecosystem 1170 fluxes due to combustion being easier to model and parameterize at large scales. A suite of eight BU methods for 1171 fossil CO₂ emissions are within the uncertainty of the NGHGI when methods are harmonized to include similar 1172 categories. The remaining differences can often be attributed to definitions, assumptions about activity data or 1173 emission factors, and the allocation of fuel types to different sectors (see Sect. 3.2 and Fig. B3). The one available TD 1174 method, a regional European inversion system (CIF-CHIMERE) using an NOx proxy to determine CO₂ fossil 1175 emissions, shows broad agreement with the BU estimates. However, this initial TD inversion is not yet capable of 1176 distinguishing the minor differences between the various BU estimates and does not yet quantify uncertainties. 1177 However, a substantial decrease in the level of uncertainty of the inverse modeling system is expected in the near-1178 term with the large-scale deployment of observation networks dedicated to detecting fossil fuel emissions (e.g., with 1179 launch of the CO₂M¹⁹ satellite mission in 2025). In the short-term, the CoCO₂ project (CoCO₂, 2022) aims to advance methodology around co-assimilation of existing CO₂ satellite data (from the OCO-2/3 instruments) and to provide 1180 1181 new analysis of the CO/FFCO2 and NOx/FFCO2 ratios in order to significantly decrease uncertainty in the fossil CO₂ 1182 estimates.

1183 The CO₂ land fluxes belong to the LULUCF sector, which is one of the most uncertain sectors in UNFCCC 1184 reporting. The IPCC guidelines prescribe methodologies that are used to estimate the CO2 fluxes in the NGHGI, but 1185 grant countries significant freedom to adopt methods appropriate to their national circumstances. When analyzing the 1186 different estimates from multiple BU sources (inventories and models) similar sources of uncertainties are observed 1187 such as: (a) differences due to input data and structural/parametric uncertainty of models (Houghton et al., 2012; 1188 Pongratz et al., 2021) and (b) differences in definitions (Pongratz et al., 2014; Grassi et al., 2018b; Petrescu et al., 1189 2020, 2021b; Grassi et al., 2021). Reducing uncertainties in LULUCF estimates is needed given the increasing 1190 importance of the sector to EU climate policy over the next decades. In contrast to the previous 2020 climate and energy package, the LULUCF sector will now formally contribute to the binding emission reduction targets of the 1191 1192 Unions 2030 climate and energy framework (EU, 2018a; 2018b). Furthermore, the European Climate Law explicitly 1193 states that LULUCF, together with all sectors of the economy, should contribute to achieving Climate neutrality within 1194 the Union by 2050 (EU, 2021b).

1195The LULUCF sector in NGHGIs is composed of six land use categories. Of these, Forest land provides the1196most important contribution to the net CO2 land flux in the EU27+UK, followed by Cropland and Grassland. HWP1197and "Land converted to settlements" also have non-negligible contributions, and changes in HWP strongly influence

¹⁹ CO₂M: Copernicus Anthropogenic Carbon Dioxide Monitoring,

https://esamultimedia.esa.int/docs/EarthObservation/CO2M_MRD_v3.0_20201001_Issued.pdf





1198 variations in decennial mean net LULUCF fluxes for the region. Of these, all except "Land converted to settlements" 1199 are represented in general ecosystem models, while Forestland, Cropland, and Grassland are simulated by sector-1200 specific process-based and data-driven models. Top-down inversions are capable of simulating net CO₂ fluxes to the 1201 atmosphere, but cannot yet attribute them between different categories.

1202 Differences in the detailed sector-specific and inversion model results (Fig. 4-9) often come from choices in 1203 the simulation setup and the type of model used: bookkeeping models, process-based DGVMs, inventory-based 1204 statistical methods, or atmospheric inversions. Results also differ based on whether fluxes are attributed to LULUCF 1205 emissions due to the cause or location of occurrence. For example, indirect fluxes on managed land are included in 1206 NGHGI and FAOSTAT, while additional sink capacity (e.g., Petrescu et al., 2021b) is included in estimates from process-based models (e.g., ORCHIDEE or TRENDY DGVMs). The use of gross land use changes fluxes (e.g., in 1207 1208 the NGHGI, bookkeeping models, and CABLE-POP) as opposed to net fluxes also likely plays an important role. We 1209 found that adjusting top-down models by emissions/removals resulting from later transport of carbon through trade 1210 and the inland water network improves the agreement with the NGHGI of the EU27+UK (Fig. 9, compared to Petrescu 1211 et al., 2021b).

1212 Observation-based BU estimates of LULUCF provide large year-to-year flux variability (Fig. 4-7, in 1213 particular for DGVMs like ORCHIDEE, CABLE-POP and the TRENDY ensemble), contrary to the NGHGI, 1214 primarily due to the effect of varying meteorology. In particular, the duration and intensity of the summer growing season can vary significantly between years (e.g., Bastos et al., 2020a; Thompson et al., 2020). In the framework of 1215 periodic NGHGI assessments, the choice of a reference period (such as 2015-2019, as used here) or the use of a 1216 1217 moving window to calculate the means may be critical to smooth out high inter-annual variability and facilitate 1218 comparisons. One can also imagine incorporating IAV into the NGHGIs through the use of annual anomalies of 1219 emission factors calculated from Tier 3 observation-based approaches (either BU or TD). TD estimates also show 1220 very large inter-annual variability (Fig. 9). Uncertainties in the inversion results are primarily due to uncertainties in atmospheric transport modeling, boundary conditions, technical simplifications and uncertainty inherent to the 1221 1222 limitation of the observation network. Currently, regional inversions (LUMIA, CSR and EUROCOM) are still under 1223 development and face different challenges from the coarser resolution global systems used here to represent regional 1224 results (GCP). Based on this work, it is difficult to claim that one or the other provides a more accurate result for the 1225 net CO₂ land fluxes across the EU27+UK, although two regional inversion ensembles (EUROCOM and CSR) 1226 dramatically reduced their uncertainties between the previous and current versions of this synthesis, with CSR showing 1227 much more overlap now with the NGHGI (Fig. 10).

Uncertainties can be reflected in space as well as in time. Fig. 11 separates mean BU and TD values for all methods into five different regions in Europe. From this figure, it's clear that some regions suffer from higher uncertainties than others. Part of this is likely linked to the spareness of atmospheric observation data for the TD estimates (e.g., Eastern Europe). Reconciling differences across aggregated EU regions may be challenging due to diverse methodologies and drivers in each country. On the other hand, the analysis of smaller regions or individual countries may represent a productive first step towards monitoring the current state of emissions as national data and experts can be used to help clarify differences across models. Country-level case studies may help inform the design





of future monitoring and verification systems (MVS) for CO₂ which aim to supply additional evidence for the
emissions levels and trends, coupling anthropogenic activities and associated emissions with the atmospheric patterns
of greenhouse gas concentrations, and perform data assimilation and modeling over a wide variety of environmental
conditions (Pinty et al., 2017).

1239 As seen in figures throughout this work, reducing uncertainties of both individual models and classes of 1240 models remains a priority. Some categories (Forestland, Cropland) produce results for multiple category-specific 1241 models which lie within the uncertainty of the NGHGI. This likely reflects relatively the use of data-driven models 1242 and the relatively high quality of data that is available due to the economic importance of these categories. On the 1243 other hand, generalized ecosystem models (the DGVMs, like ORCHIDEE and CABLE-POP) may create mean estimates which fall within uncertainties, but fall outside of NGHGI uncertainties for any given year due to the 1244 1245 sensitivity of processes in these models to rapidly changing meteorology and the necessity for these models to operate globally, including in data-poor regions for which parameterization may be impossible. Two advances in 1246 1247 characterizing uncertainty were presented here: one for the case of the NGHGI, and one for the case of the TD model 1248 CSR. Additional characterization of uncertainty both within and across models will enable more fair comparisons 1249 between methods.

1250 A more detailed analysis of LULUCF fluxes at the regional/country level is foreseen as part of projects linked to VERIFY including the RECCAP2 initiative (RECCAP2, 2022) and current and future Horizon Europe funded 1251 projects (e.g., CoCO2, EYE-CLIMA, AVENGERS, PARIS) which will highlight examples of good practice in 1252 1253 LULUCF flux monitoring amongst European countries. Sect. 3.3.6 presents a summary of uncertainties to provide 1254 insight into ground observation systems assimilated by inversions. This lays the basis of future improvements for 1255 establishing best practices on how to configure atmospheric inversions and systematically quantify uncertainties. For 1256 the overall estimation of emissions from LULUCF activities on all land types (Fig. 8), the comparison is made more 1257 challenging as results from both land use and land use changes are presented. Comparing only the "effect of land use 1258 change" (conversion) is non-trivial. A methodology for reconciling LULUCF country estimates from the FAOSTAT 1259 datasets with the NGHGIs is presented in Grassi et al. (2022a) and Grassi et al. (in prep) for the global scale.

1260 The next steps needed to improve and facilitate the reconciliation between BU and TD estimates are the same 1261 as those discussed in Petrescu et al. (2021b): 1) BU process-based models incorporating unified protocols and 1262 guidelines for uniform definitions should be able to disaggregate their estimates to facilitate comparison to NGHGI and 2006 IPCC practices (e.g., managed vs. unmanaged land, 20-year legacy for classes remaining in the same class, 1263 1264 distinction of fluxes arising solely from land use change, Grassi et al. (2022a)); 2) for sector-specific models, in 1265 particular for cropland and grassland, improving treatment of the contribution of soil organic carbon dynamics to 1266 the budget; 3) for TD estimates, using the recently developed Community Inversion Framework (Berchet et al., 2021) to better assess the different sources of uncertainties from the inversion set-ups (model transport, prior fluxes, 1267 1268 observation networks), 4) standardize methods to compare datasets with and without interannual variability, and 5) 1269 develop a clear way to report key system boundary, data, or definitional issues, as it often necessary to have deep 1270 understanding of each estimate to know how to do a like-for-like comparison.





1271 Similar to Petrescu et al. (2021b), this updated study concludes that a complete, ready-for-purpose monitoring 1272 system providing annual carbon fluxes across Europe is still under development, but data sources are beginning to 1273 show improved agreement compared to previous estimates. Therefore, significant effort must still be undertaken to 1274 reduce the uncertainty across all potential methods (i.e., structural uncertainty in the models as well as the input data 1275 supplied to the models or inventory approaches) used in such a system (e.g. Maenhout et al., 2020). Future activities 1276 in the CoCO₂ project (CoCO2, 2022) will investigate the one and five-year carbon budgets across the data-rich area 1277 of the EU27+UK and deepen the analysis for both global and regional/local (city level) estimates.

1278 Achieving the well-below 2°C temperature goal of the Paris Agreement requires, among other things, low-1279 carbon energy technologies, forest-based mitigation approaches, and engineered carbon dioxide removal (Grassi et 1280 al., 2018a; Nabuurs et al. 2017). Currently, the EU27+UK reports a sink for LULUCF and forest management will 1281 continue to be the main driver affecting the productivity of European forests for the next decades (Koehl et al., 2010), 1282 shown as well by the domination of Forestland CO₂ fluxes to the LULUCF sector in the NGHGI for the bloc. Forest 1283 management changes forest composition and structure, which affects the exchange of energy with the atmosphere 1284 (Naudts et al., 2016), and therefore the potential of mitigating climate change (Luyssaert et al., 2018; Grassi et al., 1285 2019). Meteorological extremes can also affect the efficiency of the sink (Thompson et al., 2020). The EU forest sink 1286 is projected to decrease in the near future (Vizzarri et al., 2021). Consequently, for the EU to meet its ambitious climate 1287 targets, it is necessary to maintain and even strengthen the LULUCF sink (EU, 2020). Understanding the evolution of 1288 the CO₂ land fluxes is critical to enable the EU27+UK to meet its ambitious climate goals.

1289

1290 6. Appendices

1291

1292 Appendix A: Data sources, methodology and uncertainty descriptions

Plots for all countries in Europe as well as dozens of country groups and some countries outside of Europe are availablefollowing a simple registration (VERIFY Synthesis Plots, 2022).

1295

1296 VERIFY project

VERIFY's primary aim is to develop scientifically robust methods to assess the accuracy and potential biases
 in national inventories reported by the parties through an independent pre-operational framework. The main concept
 is to provide observation-based estimates of anthropogenic and natural GHG emissions and sinks as well as associated
 uncertainties. The proposed approach is based on the integration of atmospheric measurements, improved emission
 inventories, ecosystem data, and satellite observations, and on an understanding of processes controlling GHG fluxes
 (ecosystem models, GHG emission models).

1303 Two complementary approaches relying on observational data-streams were combined in VERIFY to1304 quantify GHG fluxes:

1305 1) atmospheric GHG concentrations from satellites and ground-based networks (top-down atmospheric inversion1306 models) and





- 1307 2) bottom-up activity data (e.g., fuel use and emission factors) and ecosystem measurements (bottom-up models).
- **1308** For CO₂, a specific effort was made to separate fossil fuel emissions from ecosystem fluxes.
- 1309

1310 The objectives of VERIFY were:

- 1311 Objective 1. Integrate the efforts between the research community, national inventory compilers, operational centers
- 1312 in Europe, and international organizations towards the definition of future international standards for the verification
- 1313 of GHG emissions and sinks based on independent observation.
- 1314 **Objective 2**. Enhance the current observation and modeling ability to accurately and transparently quantify the sinks
- 1315 and sources of GHGs in the land-use sector for the tracking of land-based mitigation activities.
- 1316 Objective 3. Develop new research approaches to monitor anthropogenic GHG emissions in support of the EU
- 1317 commitment to reduce its GHG emissions by 40 % by 2030 compared to the year 1990.
- 1318 Objective 4. Produce periodic scientific syntheses of observation-based GHG balance of EU countries and practical
- 1319 policy-oriented assessments of GHG emission trends, and apply these methodologies to other countries.
- 1320
- 1321 For more information on the project team and products/results please visit the VERIFY website (VERIFY, 2022).
- 1322

Country name – geographical Europe	BU-ISO3	Aggregation from TD-ISO3
Luxembourg	LUX	
Belgium	BEL	BENELUX
Netherlands	NLD	BNL
Bulgaria	BGR	BGR
Switzerland	CHE	
Lichtenstein	LIE	CHL
Czech Republic	CZE	Former Czechoslovakia
Slovakia	SVK	CSK
Austria	AUT	AUT
Slovenia	SVN	North Adriatic countries
Croatia	HRV	NAC
Romania	ROU	ROU
Hungary	HUN	HUN
Estonia	EST	

1323Table A1: Country grouping used for comparison purposes between BU and TD emissions as reported for the1324country- and regional-level synthesis plots available through the VERIFY web portal.





Lithuania	LTU	Baltic countries
Latvia	LVA	BLT
Norway	NOR	NOR
Denmark	DNK	
Sweden	SWE	
Finland	FIN	DSF
Iceland	ISL	ISL
Malta	MLT	MLT
Cyprus	СҮР	СҮР
France (Corsica incl.)	FRA	FRA
Monaco	МСО	
Andorra	AND	
Italy (Sardinia, Vatican incl.)	ITA	ITA
San Marino	SMR	
United Kingdom (Great Britain + N Ireland)	GBR	UK
Isle of Man	IMN	
Iceland		
Ireland	IRL	IRL
Germany	DEU	DEU
Spain	ESP	IBERIA
Portugal	PRT	IBE
Greece	GRC	GRC
Russia (European part)	RUS European	
Georgia	GEO	RUS European+GEO
Russian Federation	RUS	RUS
Poland	POL	POL
Turkey	TUR	TUR
EU27+UK (Austria, Belgium, Bulgaria, Cyprus, Czech Republic, Germany, Denmark, Spain, Estonia, Finland, France, Greece, Croatia, Hungary, Ireland, Italy, Lithuania, Latvia, Luxembourg,	AUT, BEL, BGR, CYP, CZE, DEU, DNK, ESP, EST, FIN, FRA, GRC, HRV, HUN, IRL.	E28





Malta, Netherlands, Poland, Portugal, Romania, Slovakia, Slovenia, Sweden, United Kingdom)	ITA, LTU, LVA, LUX, MLT, NDL, POL, PRT, ROU, SVN, SVK, SWE, GBR	
Western Europe (Belgium, France, United Kingdom, Ireland, Luxembourg, Netherlands)	BEL, FRA, UK, IRL, LUX, NDL	WEE
Central Europe (Austria, Switzerland, Czech Republic, Germany, Hungary, Poland, Slovakia)	AUT, CHE, CZE, DEU, HUN, POL, SVK	CEE
Northern Europe (Denmark, Estonia, Finland, Lithuania, Latvia, Norway, Sweden)	DNK, EST, FIN, LTU, LVA, NOR, SWE	NOE
South-Western Europe (Spain, Italy, Malta, Portugal)	ESP, ITA, MLT, PRT	SWN
South-Eastern Europe (all) (Albania, Bulgaria, Bosnia and Herzegovina, Cyprus, Georgia, Greece, Croatia, Macedonia, the former Yugoslav, Montenegro, Romania, Serbia, Slovenia, Turkey)	ALB, BGR, BIH, CYP, GEO, GRC, HRV, MKD, MNE, ROU, SRB, SVN, TUR	SEE
South-Eastern Europe (Albania, Bosnia and Herzegovina, Macedonia, the former Yugoslav, Georgia, Turkey, Montenegro, Serbia)	ALB, BIH, MKD, MNE, SRB, GEO, TUR	SEA
South-Eastern Europe (EU) (Bulgaria, Cyprus, Greece, Croatia, Romania, Slovenia)	BGR, CYP, GRC, HRV, ROU, SVN	SEZ
Southern Europe (all) (SOE) (Albania, Bulgaria, Bosnia and Herzegovina, Cyprus, Georgia, Greece, Croatia, Macedonia, the former Yugoslav, Montenegro, Romania, Serbia, Slovenia, Turkey, Italy, Malta, Portugal, Spain)	ALB, BGR, BIH, CYP, GEO, GRC, HRV, MKD, MNE, ROU, SRB, SVN, TUR, ITA, MLT, PRT, ESP	SOE
Southern Europe (SOY) Albania, Bosnia and Herzegovina, Georgia, Macedonia, the former Yugoslav, Montenegro, Serbia, Turkey)	ALB, BIH, GEO, MKD, MNE, SRB, TUR,	SOY
Southern Europe (EU) (SOZ) (Bulgaria, Cyprus, Greece, Croatia, Romania, Slovenia, Italy, Malta, Portugal, Spain)	BGR, CYP, GRC, HRV, ROU, SVN, ITA, MLT, PRT, ESP	SOZ
Eastern Europe (Belarus, Moldova, Republic of, <i>Russian Federation</i> , Ukraine)	BLR, MDA, <i>RUS,</i> UKR	EAE
EU-15 (Austria, Belgium, Germany, Denmark, Spain, Finland, France, United Kingdom, Greece, Ireland, Italy, Luxembourg, Netherlands, Portugal, Sweden)	AUT, BEL, DEU, DNK, ESP, FIN, FRA, GBR, GRC, IRL, ITA, LUX, NDL, PRT, SWE	E15
EU-27 (Austria, Belgium, Bulgaria, Cyprus, Czech Republic, Germany, Denmark, Spain, Estonia, Finland, France, Greece, Croatia, Hungary, Ireland, Italy, Lithuania, Latvia, Luxembourg,	AUT, BEL, BGR, CYP, CZE, DEU, DNK, ESP, EST, FIN, FRA, GRC, HRV,	E27





Malta, Netherlands, Poland, Portugal, Romania, Slovakia, Slovenia, Sweden)	HUN, IRL. ITA, LTU, LVA, LUX, MLT, NDL, POL, PRT, ROU, SVN, SVK, SWE	
All Europe (Aaland Islands, Albania, Andorra, Austria, Belgium, Bulgaria, Bosnia and Herzegovina, Belarus, Switzerland, Cyprus, Czech Republic, Germany, Denmark, Spain, Estonia, Finland, France, Faroe Islands, United Kingdom, Guernsey, Greece, Croatia, Hungary, Isle of Man, Ireland, Iceland, Italy, Jersey, Liechtenstein, Lithuania, Luxembourg, Latvia, Moldova, Republic of, Macedonia, the former Yugoslav, Malta, Montenegro, Netherlands, Norway, Poland, Portugal, Romania, Russian Federation, Svalbard and Jan Mayen, San Marino, Serbia, Slovakia, Slovenia, Sweden, Turkey, Ukraine)	ALA, ALB, AND, AUT, BEL, BGR, BIH, BLR, CHE, CYP, CZE, DEU, DNK, ESP, EST, FIN, FRA, FRO, GBR, GGY, GRC, HRV, HUN, IMN, IRL, ISL, ITA, JEY, LIE, LTU, LUX, LVA, MDA, MKD, MLT, MNE, NDL, NOR, POL, PRT, ROU, RUS, SJM, SMR, SRB, SVK, SVN, SWE, TUR, UKR	EUR

1325 *countries highlighted in *italic* are not discussed in the current 2021 synthesis mostly because unavailability of UNFCCC NGHGI reports (non-

1326 Annex I countries²⁰) but are present on the web-portal (VERIFY Synthesis Plots, 2022). Results for Annex I countries (NOR, CHE, ISL) and

1327 Eastern European countries (EAE) are represented in Fig. 11.

²⁰Non-Annex I countries are mostly developing countries. The reporting to UNFCCC is implemented through national communications (NCs) and biennial update reports (BURs): <u>https://unfccc.int/national-reports-from-non-annex-i-parties</u>





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1330 *Table A2: Methodological changes (in bold) of the current study with respect to Petrescu et al. (2020), Petrescu et al.*

1331 (2021b) and an internal VERIFY update (v2020); n/a cells mean that there is no data available.

1332

Publication year	Bottom-up a estimates (fossil CO ₂)	inthropogenic CC	\mathcal{D}_2	Top-down fossil CO ₂ estimates	Bottom-up natural CO ₂ (NBP) emissions/removals (land CO ₂)			Top-down land emissions	d CO ₂	Uncertainty and other changes
	Inventorie s	Global databases	Emissio n models		Inventories	Emission models	Global Databases	Regional models	Global models	
Petrescu et al. (2020) AFOLU bottom-up synthesis	n/a	n/a	n/a	n/a	National emissions from UNFCCC (2018) 1990-2016	CBM Forest land (2000, 2005, 2010 and 2015)	FAOSTAT Timeseries Remaining and conversions 1990-2016	n/a	n/a	UNFCCC (2018) uncertainty estimates for 2016 (error propagation 95 % interval method)
					LULUCF Forest land, -	EFISCEN Forest land				
					EU28 data for five years (1995, 2000, 2005, 2010 and 2015)	(1995, 2000, 2005, 2010 and 2015)				
					Cropland and Grassland	BLUE				
					(1990, 2005, 2010 and 2016)	All land uses				
					All land uses					
					EU28 timeseries	H&N				
					1990-2016	All land uses 1990-2015				
						DGVMs (TRENDY v6)				
						All land uses				
~						1990-2017				
Petrescu et al., 2021b	National emissions from UNFCCC	EDGAR v5.0 BP	n/a	IAP RAS fast- track inversion	National emissions from UNFCCC	CBM Forest land	FAOSTAT	CSR 2006-2018	GCP 2019 inversions	UNFCCC (2019) uncertainty estimates for
					(2019)	1 imeseries			2000-2018	2016 (error propagation





(2019)	EIA	2014	1990-2017	1990-2015	Remaining	EUROCOM	95 % interval
CRFs	CDIAC	(EU11+CHE)	$EU27 \pm UK$		and .	2006 2015	method)
2014	CDIAC		timeseries		conversions	2000-2015	
	IEA			EFISCEN	1990-2017		
	GCP		of Forest	Forest land			For model
All			Cropland and				reported as
anthropog	CEDS		Grassland	timeseries			variability in
enic (excl.				2005-2018			extremes
) sectors,	2014						(min/max)
timeseries	2014 estimates		Regional				
1990-	split by fuel		EU27 + UK totals (incl	CO ₂			
2015	type		NOR, CHE,	emissions			
			UKR, MLD	waters			
			and BLR)				
	EDGAR						
	v3.0			ORCHIDEE			
	All			Forest			
	anthropogeni c sectors			cropland and			
	timeseries			grassland and			
	1000 2015			all land uses			
	1990-2015			1990-2018			
				ECOSSE			
				Cropland and			
				grassland			
				1000 2018			
				1990-2018			
				EPIC-IIASA			
				Cropland			
				1990-2018			
				BLUE			
				All land uses			
				1000 2018			
				1990-2018			
				H&N			
				All land uses			
				1990-2015			
				DGVMs (TRENDY			
				(1 KEND Y)			
				,			
				All land uses			





			1990-2018		
					l





This study	National	EDGAR	n/a	CIF-	National	CBM	FAOSTAT	CarboScope	GCP 2021	UNFCCC
	emissions	v6.0		CHIMERE	emissions	TT: - 1 0		Reg	inversions	(2021)
	from	DD		fast-track	from	Historical flux	timeseries	2006 2020		uncertainty
	UNFCCC	вр		inversion	UNFCCC	from Forest	Remaining	2006-2020	2010-2020	estimates for
	(2021) CDEs	EIA		2005-2020	(2021)	land	and			2019 (error
	CKFS			(EU27+UK)	1990- 2019	remaining	conversions			95 % interval
	2017	CDIAC				forest land		EUROCOM		method)
		IFA			EU27 + UK		1990- 2019	2009-2018		inteniou)
		ILA			of Forest land	2000-2015		2009-2010		
	A11	GCP			of Porest fand	and new		LUMIA		E an ana dal
	anthropog	GEDG				2017-2020 ostimato				ensembles
	enic (excl.	CEDS				comate		2006-2020		reported as
	LULUCF	PRIMAP-								the annual
) sectors,	hist 2.3								extremes
	timeseries	UNFCCC				EFISCEN-		CIF-		(min/max)
	1990-	NGHGI				STACE (undated		CHIMERE		
	2019	2021				(updated model)		2005-2020		
	-015					Forest land		2003 2020		
		2017				timeseries				
		estimates				2005-2020				
		split by fuel				For 15 EU				
		type				countries				
		EDGAR				OBCUIDEE				
		v6.0				ORCHIDEE				
		A 11				Forest,				
		All				cropland and				
		c sectors				grassland and				
		timeseries				all land uses,				
						model				
		1990-2018				updated				
						1990-2020				
						CABLE-POP				
						1990-2020				
						BLUE-				
						VERIFY and				
						BLUE-GCP)				
						All land uses				
						1000 2010				
						1990-2019,				
						1550-2020				
						THO N				
						H&N CCP2021				
						All land uses				
						. In fund uses				
						1990 -2020				
						DGVMs				
						(TRENDY				
						v10)				





			Forest,		
			cropland and		
			grassiand and		
			an land uses		
			1990-2020		
			ECOSSE		
			Cropland		
			Cropiand		
			1990 -2020		
			Grassland		
			1990-2018		
			EPIC-IIASA		
			Cropland		
			1000 2020		
			1990-2020		
			Grassland		
			1990-2020		
			1770-2020		

1333

1334

- 1335 A1: Fossil CO₂ emissions
- 1336 Bottom-up emission estimates
- 1337 For further details of all datasets, see Andrew (2020).

1338

1339 UNFCCC NGHGI (2021)

Annex I NGHGIs should follow principles of transparency, accuracy, consistency, completeness and comparability (TACCC) under the guidance of the UNFCCC (UNFCCC, 2014) and as mentioned above, shall be completed following the 2006 IPCC guidelines (IPCC, 2006). In addition, the IPCC 2019 Refinement (IPCC, 2019), which may be used to complement the 2006 IPCC guidelines, has updated sectors with additional emission sources and provides guidance on the use of atmospheric data for independent verification of GHG inventories.

Both approaches (BU and TD) provide useful insights on emissions from two different points of view. First, as outlined in Volume 1, Chapter 6 of the 2019 IPCC Refinement (IPCC, 2019), TD approaches act as an additional quality check for BU and NGHGI approaches, and facilitate a deeper understanding of the processes driving changes in different elements of GHG budgets. Second, while independent BU methods do not follow prescribed standards like the IPCC Guidelines, they do provide complementary information based on alternative input data at varying temporal, spatial, and sectoral resolution. This complementary information helps build trust in country GHG estimates,





which form the basis of national climate mitigation policies. Additionally, BU estimates are needed as input for TD estimates. As there is no formal guideline to estimate uncertainties in TD or BU approaches, uncertainties are usually assessed from the spread of different estimates within the same approach, though some groups or institutions report uncertainties for their individual estimates using a variety of methods, for instance, by performing Monte Carlo sensitivity simulation by varying input data parameters. However, this can be logistically and computationally difficult when dealing with complex process-based models.

1357 Despite the important insights gained from complementary BU and TD emission estimates, it should be noted 1358 that comparisons with the NGHGI are not always straightforward. BU estimates often share common methodology 1359 and input data, and through harmonization, structural differences between BU estimates and NGHGIs can be 1360 interpreted. However, the use of common input data restricts the independence between the datasets and, from a 1361 verification perspective, may limit the conclusions drawn from the comparisons. On the other hand, TD estimates are 1362 constrained by independent atmospheric observations and can serve as an additional, nearly independent quality check 1363 for NGHGIs. Nonetheless, structural differences between NGHGIs (what sources and sinks are included, and where 1364 and when emissions/removals occur) and the actual fluxes of GHGs to the atmosphere must be taken into account 1365 during comparison of estimates. While NGHGIs go through a central QA/QC review process, the UNFCCC reporting 1366 requirements do not mandate large-scale observation-derived verification. Nevertheless, the individual countries may 1367 use atmospheric data and inverse modeling within their data quality control, quality assurance and verification processes, with expanded and updated guidance provided in chapter 6 of the 2019 Refinement of IPCC 2006 1368 Guidelines (IPCC, 2019). So far, only a few countries (e.g. Switzerland, UK, New Zealand and Australia) have used 1369 1370 atmospheric observations to constrain national emissions and documented these verification activities in their national 1371 inventory reports (Bergamaschi et al., 2018), and none do so for CO₂.

1372 Under the UNFCCC convention and its Kyoto Protocol, national greenhouse gas (GHG) inventories are the 1373 most important source of information to track progress and assess climate protection measures by countries. In order 1374 to build mutual trust in the reliability of GHG emission information provided, national GHG inventories are subject 1375 to standardized reporting requirements, which have been continuously developed by the Conference of the Parties 1376 $(COP)^{21}$. The calculation methods for the estimation of greenhouse gasses in the respective sectors is determined by 1377 the methods provided by the 2006 IPCC Guidelines for National Greenhouse Gas Inventories (IPCC, 2006). These 1378 Guidelines provide detailed methodological descriptions to estimate emissions and removals, as well as 1379 recommendations to collect the activity data needed. As a general overall requirement, the UNFCCC reporting 1380 guidelines stipulate that reporting under the Convention and the Kyoto Protocol must follow the five key principles 1381 of transparency, accuracy, completeness, consistency and comparability (TACCC).

1382The reporting under UNFCCC shall meet the TACCC principles. The three main GHGs are reported in1383timeseries from 1990 up to two years before the due date of the reporting. The reporting is strictly source category1384based and is done under the Common Reporting Format tables (CRF), downloadable from the UNFCCC official1385submission portal: https://unfccc.int/ghg-inventories-annex-i-parties/2021.

²¹ The last revision has been made by COP 19 in 2013 (UNFCCC, 2013)



1386The UNFCCC NGHGI CO2 emissions/removals include estimates from five key sectors for the EU27+UK:13871 Energy, 2 Industrial processes and product use (IPPU), 3 Agriculture, 4 LULUCF and 5 Waste. The tiers method a1388country applies depends on the national circumstances and the individual conditions of the land, which explains the1389variability of uncertainties among the sector itself as well as among EU countries. This annual published dataset1390includes all CO2 emissions sources for those countries, and for most countries for the period 1990 to t-2. Some eastern1391European countries' submissions began in the 1980s.

1392

1393 NGHGI uncertainties

1394 The presented uncertainties in the reported emissions of the individual countries and the EU27+UK bloc 1395 were calculated by using the methods and data used to compile the official GHG emission uncertainties that are 1396 reported by the EU under the UNFCCC (NIRs, 2022). The EU uncertainty analysis reported in the bloc's National 1397 Inventory Report (NIR) is based on country-level, Approach 1 uncertainty estimates (IPCC, 2006, Vol. 1, Chap. 3) 1398 that are reported by EU Member States, Iceland and United Kingdom under Article 7(1)(p) of EU (2013). These 1399 country-level uncertainty estimates are typically reported at beginning of a submission cycle and are not always 1400 revised with updated CRF submissions later in the submission cycle. Furthermore, the compiled uncertainties of some 1401 countries are incomplete (e.g., uncertainties not estimated for LULUCF and/or indirect CO2 emissions, certain 1402 subsector emissions are confidential) and the sector and gas resolution at which uncertainties are provided varies 1403 between the countries. The EU inventory team therefore implements a procedure to harmonize and gap-fill these 1404 uncertainty estimates. A processing routine reads the individual country uncertainty files that are pre-formatted 1405 manually to assign consistent sector and gas labels to the respective estimates of emissions/removals and uncertainties. 1406 The uncertainty values are then aggregated to a common sector resolution, at which the emissions and removals 1407 reported in the uncertainty tables of the countries are then replaced with the respective values from the final CRF 1408 tables of the countries. Due to the issue of incompleteness mentioned above, the country-level data are then screened 1409 to identify residual GHG emissions and removals for which no uncertainty estimates have been provided. Where 1410 sectors are partially complete, the residual net emission is quantified in CO₂ equivalents and incorporated. An 1411 uncertainty is then estimated, by calculating the overall sector uncertainty of the sources and sinks that were included 1412 in that country's reported uncertainties estimates and assigning this percentage average to the residual net emission. 1413 In cases where for certain sectors no uncertainties have been provided at all (e.g., indirect CO2 emissions, LULUCF), 1414 an average (median) sector uncertainty in percent is calculated from all the countries for which complete sectoral 1415 emissions and uncertainties were reported, and this average uncertainty is assigned to the country's sector GHG total 1416 reported in its final CRF tables.

1417The country-level uncertainties presented in this paper, have been compiled using this same processing1418routine and using the uncertainties and CRF data reported by the countries in the 2021 submission. However, here the1419method has been expanded to gap-fill at the individual greenhouse gas level (CO₂ emissions and removals only) rather1420than at the aggregate GHG level. Furthermore, the expanded method here assigns the sub-sectoral uncertainties to the1421emissions and removals of the entire timeseries (1990-2019), rather than just the base year and latest year of the1422respective timeseries. This allows uncertainties to be sensitive to the sub-sectoral contributions to sectoral and national





1423 total emissions, which of course change over time. For each year of the timeseries, uncertainties in the total and sectoral CO2 emissions are calculated using Gaussian error propagation, by summing the respective sub-sectoral 1424 1425 uncertainties (expressed in kt CO2) in quadrature and assuming no error correlation. In contrast, for the EU27+UK 1426 bloc, uncertainties in the total and sectoral CO₂ emissions were calculated to take into account error correlations 1427 between the respective country estimates at the subsector level. This was done by applying the same methods and 1428 assumptions described in the 2022 EU NIR (UNFCCC NIR, 2022). The subsector resolution applied for gap-filling 1429 allows the routine to access respective data on emission factors from CRF Table Summary 3 and apply correlation 1430 coefficients (r) when aggregating the uncertainties. For a given subsector, it is assumed that the errors of countries 1431 using default factors are completely correlated (r = 1), while errors of countries using country-specific factors are assumed uncorrelated (r = 0). For countries using a mix of default and country-specific factors at the given subsector 1432 1433 level, it is assumed that these errors are partially correlated (r = 0.5) with one another and with the errors of countries 1434 using the default factors only.

1435 Based on these correlation assumptions, the routine then aggregates CO₂ emissions/removals and 1436 uncertainties for the specified subsector resolution at the EU27+UK level. Uncertainties at sector total level are then 1437 aggregated from the subsector estimates assuming no correlation between subsectors. However, for countries reporting 1438 very coarse resolution estimates (e.g., total sector CO₂ emissions/removals) or where the sector has been partially or 1439 completely gap-filled, it is assumed that these uncertainties are partially correlated (r = 0.5) with one another and with 1440 the other reported subsector level estimates. Level uncertainties on the total EU27+UK CO2 emissions and removals 1441 (with and without LULUCF) are then aggregated from the sector estimates assuming no error correlation between 1442 sectors.

1443Note that the above procedure does not apply to LULUCF categories (FL, CL, and GL). Estimates for these1444values were taken directly from the EU NIR (2021 without gap-filling or consideration of correlations. As the values1445are given for only one single year, this value is applied uniformly across the whole timeseries.

1446 EDGAR v6.0

1447The first edition of the Emissions Database for Global Atmospheric Research was published in 1995. The1448dataset now includes almost all sources of fossil CO2 emissions, is updated annually, and reports data for 1970 to year1449n-1. Estimates for v6.0 are provided by sector. Emissions are estimated fully based on statistical data from 1970 till14502018 https://data.jrc.ec.europa.eu/dataset/97a67d67-c62e-4826-b873-9d972c4f670b.

Uncertainties: EDGAR uses emission factors (EFs) and activity data (AD) to estimate emissions. Both EFs and AD are uncertain to some degree, and when combined, their uncertainties need to be combined too. To estimate EDGAR's uncertainties (stemming from lack of knowledge of the true value of the EF and AD), the methodology devised by IPCC (2006, Chapter 3) is adopted, that is the overall uncertainty is the square root of the sum of squares of the uncertainty of the EF and AD (uncertainty of the product of two variables). A log-normal probability distribution function is assumed in order to avoid negative values, and uncertainties are reported as the 95 % confidence interval according to IPCC (2006, chapter 3, equation 3.7). For emission uncertainty in the range 50 % to 230 % a correction





1458	factor is adopted as suggested by Frey et al. (2003) and IPCC (2006, chapter 3, equation 3.4). Uncertainties are
1459	published in Solazzo et al. (2021).
1460	BP
1461	BP releases its Statistical Review of World Energy annually in June, the first report being published in 1952.
1462	Primarily an energy dataset, BP also includes estimates of fossil-fuel CO ₂ emissions derived from its energy data (BP
1463	2011, 2017). The emissions estimates are totals for each country starting in 1965 to year n-1.
1464	CDIAC
1465	The original Carbon Dioxide Information Analysis Center included a fossil CO2 emissions dataset that was
1466	long known as CDIAC. This dataset is now produced at Appalachian State University, and has been renamed CDIAC-
1467	FF (CDIAC, 2022). It includes emissions from fossil fuels and cement production from 1751 to year n-3. Fossil-fuel
1468	emissions are derived from UN energy statistics, and cement emissions from USGS production data.
1469	EIA
1470	The US Energy Information Administration publishes international energy statistics and from these derives
1471	estimates of energy combustion CO2 emissions. Data are currently available for the period 1980-2016.
1472	IEA
1473	The International Energy Agency publishes international energy statistics and from these derives estimates
1474	of energy combustion CO_2 emissions including from the use of coal in the iron and steel industry. Emissions estimates
1475	start in 1960 for OECD members and 1971 for non-members, and run through n-1 for OECD members' totals, and
1476	year n-2 for members' details and non-members. Estimates are available by sector for a fee.
1477	GCP
1478	The Global Carbon Project includes estimates of fossil CO2 emissions in its annual Global Carbon Budget
1479	publication. These include emissions from fossil fuels and cement production for the period 1750 to year n-1.
1480	CEDS
1481	The Community Emissions Data System has included estimates of fossil CO ₂ emissions since 2018, with an
1482	irregular update cycle (CEDS, 2022). Energy data are directly from IEA, but emissions are scaled to higher-priority
1483	sources, including national inventories. Almost all emissions sources are included and estimates are published for the
1484	period 1750 to year n-1. Estimates are provided by sector.
1485	PRIMAPv2.2
1486	The PRIMAP-hist dataset combines several published datasets to create a comprehensive set of greenhouse
1487	gas emission pathways for every country and Kyoto gas, covering the years 1850 to 2018, and all UNECCC (United
1488	Nations Framework Convention on Climate Change) member states as well as most non-UNFCCC territories. The
1489	data resolves the main IPCC (Intergovernmental Panel on Climate Change) 2006 categories. For CO ₂ CH ₄ and N ₂ O
1490	subsector data for Energy, Industrial Processes and Product Use (IPPU), and Agriculture is available. Due to data
1491	availability and methodological issues, version 2.2 of the PRIMAP-hist dataset does not include emissions from Land
1.01	a substruction of the first of

1492 Use, Land-Use Change, and Forestry (LULUCF). More info at <u>https://zenodo.org/record/4479172#.YUsc6p0zbIU</u>.







1493 © VERIFY Project
 1494 Figure A1: Comparison of EU27+UK fossil CO₂ emissions from multiple inventory datasets; Identical to Fig. 2,

1495 except that no system boundaries harmonization has been done. CDIAC does not report emissions prior to 1992 for

1496 *former-Soviet Union countries. CRF: UNFCCC NGHGI from the Common Reporting Format tables.*

1497

1498

1499 Top-down CO₂ emission estimates

1500 CIF-CHIMERE - fossil CO₂ emission inversion

1501 CIF-CHIMERE is used for both CO₂ land and CO₂ fossil emission estimates, and this section only describes
 1502 the CO₂ fossil estimates. The product is explained in more detail by Fortems-Cheiney and Broquet, 2021.

1503 Results from previous atmospheric inversions of the European fossil CO2 emissions indicated that there were 1504 much larger uncertainties associated with the assimilation of CO data than with that of NO2 data for such a purpose 1505 (Konovalov et al, 2016; Konovalov and Llova, 2018). In this context, we have developed an atmospheric inversion 1506 configuration quantifying monthly to annual budgets of the national emissions of fossil CO₂ in Europe based on the 1507 assimilation of the long-term series of NO2 spaceborne observations; the Community Inversion Framework (CIF); the 1508 CHIMERE regional chemistry transport model (CTM); corrections to the TNO-GHGco-v3 inventory of NOx 1509 anthropogenic emissions at 0.5° horizontal resolution; and the conversion of NOx anthropogenic emission estimates 1510 into CO2 fossil emission estimates. For the first time, to our knowledge, variational regional inversions have been 1511 performed to estimate the European CO₂ fossil emissions using NOx emissions from OMI satellite observations. 1512 Particular attention is paid in the analysis assessing the consistency between the fossil CO2 emissions estimates from 1513 our processing chain with the fossil CO2 emission budgets provided by the TNO-GHGco-v3 inventory based on the





emissions reported by countries to UNFCCC, which are assumed to be accurate in Europe. The algorithm first
optimizes NOx emissions and then assumes a fixed ratio of NOx to fossil CO₂ emissions. However, long-term plans
include the simultaneous inversion of all three gasses (CO₂, NO₂, and CO).

1517 The analysis is conducted over the period 2005 to 2020. CHIMERE is run over a 0.5°×0.5° regular grid and 1518 17 vertical layers, from the surface to 200hPa, with 8 layers within the first two kilometers. The domain includes 101 (longitude) x 85 (latitude) grid-cells (15.25°W-35.75°E; 31.75°N-74.25°N) and covers Europe. CHIMERE is driven 1519 1520 by the European Centre for Medium-Range Weather Forecasts (ECMWF) meteorological forecast (Owens and 1521 Hewson, 2018). The chemical scheme used in CHIMERE is MELCHIOR-2, with more than 100 reactions (Lattuati, 1522 1997; CHIMERE 2017), including 24 for inorganic chemistry. Climatological values from the LMDZ-INCA global 1523 model (Szopa et al., 2008) are used to prescribe concentrations at the lateral and top boundaries and the initial 1524 atmospheric composition in the domain. Considering the short NO2 lifetime, we do not consider its import from outside 1525 the domain: its boundary conditions are set to zero. Nevertheless, we take into account peroxyacetyl nitrate (PAN) 1526 and the associated NOx reservoir for the large-scale transport of NOx.

- Several critical aspects of this workflow need to be highlighted: (i) Fortems-Cheiney and Broquet (2021) have not yet reported estimates of the uncertainty in the fossil CO₂ emissions (this requires the derivation of the uncertainties in the NOx emission inversions and in the NOx-to-FFCO2 emission conversion), and (ii) the fossil CO₂ emission budgets provided by the TNO-GHGco-v3 inventory are based on the emissions reported by countries to UNFCCC, which are assumed to be accurate in Europe, and therefore the NOx inversion prior estimate is consistent with the inventory estimates (with respect to the NOx-to-FFCO2 emission conversion used to infer fossil CO₂ emissions from the NOx inversions).
- 1534 Uncertainty: There is no uncertainty estimate currently available for this product.
- 1535

1536 A2: Land CO₂ emissions/removals

1537 Bottom-up CO₂ estimates

1538 UNFCCC NGHGI 2021 - LULUCF

Under the convention and its Kyoto Protocol, national greenhouse gas (GHG) inventories are the most important source of information to track progress and assess climate protection measures by countries. In order to build mutual trust in the reliability of GHG emission information provided, national GHG inventories are subject to standardized reporting requirements, which have been continuously developed by the Conference of the Parties (COP)²². The calculation methods for the estimation of greenhouse gasses in the respective sectors is determined by the methods provided by the 2006 IPCC Guidelines for National Greenhouse Gas Inventories (IPCC, 2006). They provide detailed methodological descriptions to estimate emissions and removals, as well as recommendations to

²²The last revision has been made by COP 19 in 2013 (UNFCCC, 2013)





collect the activity data needed. As a general overall requirement, the UNFCCC reporting guidelines stipulate that
reporting under the Convention and the Kyoto Protocol must follow the five key principles of transparency, accuracy,
completeness, consistency and comparability (TACCC).

1549 The reporting under UNFCCC shall meet the TACCC principles. The three main GHGs are reported in 1550 timeseries from 1990 up to two years before the due date of the reporting. The reporting is strictly source category 1551 based and is done under the Common Reporting Format tables (CRF), downloadable from the UNFCCC official 1552 submission portal: <u>https://unfccc.int/ghg-inventories-annex-i-parties/2021</u>.

1553 For the biogenic CO₂ emissions from sector 4 LULUCF, methods for the estimation of CO₂ removals differ 1554 enormously among countries and land use categories. Each country uses its own country specific method which takes into account specific national circumstances (as long as they are in accordance with the 2006 IPCC guidelines), as 1555 1556 well as IPCC default values, which are usually more conservative and result in higher uncertainties. The EU GHG 1557 inventory underlies the assumption that the individual use of national country specific methods leads to more accurate 1558 GHG estimates than the implementation of a single EU wide approach (UNFCCC, 2018b). Key categories for the 1559 EU27 are 4.A.1 Forest Land: Land Use CO2, 4.A.2. Forest Land: Land Use CO2, 4.B.1 Cropland Land Use CO2, 4.B.2 1560 Cropland Land Use CO2, 4.C.1 Grassland Land Use CO2, 4.C.2 Grassland Land Use CO2, 4.D.1 Wetlands Land Use 1561 CO2, 4.E.2 Settlements Land Use CO2, and 4.G Harvested Wood Production Wood product CO2. The tiered method 1562 a country applies depends on the national circumstances and the individual conditions of the land, which explains the 1563 variability of uncertainties among the sector itself as well as among EU countries.

Table A3 shows the mean values of all LULUCF categories for the EU27+UK NGHGI (2021). The contribution is calculated as the percentage of the sum of the absolute values of all the categories, in order to account for differing signs.

1567 Table A3: LULUCF categories for the EU27+UK NGHGI (2021)

Category	Mean value for 1990-2020 [Tg C]	Contribution to gross LULUCF flux [%]
Forest land remaining forest land	-107	56.0
Land converted to forest land	-13.0	6.80
Cropland remaining cropland	8.45	4.41
Land converted to cropland	14.0	7.33
Grassland remaining grassland	11.8	6.16
Land converted to grassland	-8.22	4.23
Wetlands remaining wetlands	2.89	1.51
Land converted to wetlands	1.09	0.567
Settlements remaining settlements	1.42	0.744
Land converted to settlements	11.8	6.15





Other land remaining other land	N/A	N/A
Land converted to other land	0.135	0.0706
Harvested wood products	-11.5	5.99

1568 1569



1570

1571 Figure A2: The gains, losses, and total HWP pools from the Common Reporting Format tables for the European

Union (Convention), which covers the EU27+UK. Dashed lines show the averages for 1990-1999, 2000-2009, and
2010-2019 for easy comparison with Fig. 3.

1574

1575 Uncertainty: Methodology for the NGHGI UNFCCC submissions are based on Chapter 3 of 2006 IPCC Guidelines1576 for National Greenhouse Gas Inventories and is the same as described in Appendix A1.

1577

1578 ORCHIDEE

1579 ORCHIDEE is a general ecosystem model designed to be coupled to an atmospheric model in the context of 1580 modeling the entire Earth system. As such, ORCHIDEE calculates its prognostic variables (i.e., a multitude of C, 1581 H2O and energy fluxes) from the following environmental drivers: air temperature, wind speed, solar radiation, air 1582 humidity, precipitation and atmospheric CO2 concentration. As the run progresses, vegetation grows on each pixel, 1583 divided into fifteen generic types (e.g., broadleaf temperate forests, C3 crops), which cycle carbon between the soil, 1584 land surface, and atmosphere, through such processes such as photosynthesis, litter fall, and decay. Limited human 1585 activities are included through the form of generic wood and crop harvests, which remove aboveground biomass on 1586 an annual basis. The version reported here, ORCHIDEE-N v3, includes a dynamic nitrogen cycle coupled to the





vegetation carbon cycle which results in, among other things, limitations on photosynthesis in nitrogen-poorenvironments (Vuichard et al., 2019)

1589 Among other environmental indicators, ORCHIDEE simulates positive and negative CO2 emissions from 1590 plant uptake, soil decomposition, and harvests across forests, grasslands, and croplands. Activity data is based on land 1591 use and land cover maps. For VERIFY, pixel land cover/land use fractions were based on a combination of the land 1592 use map LUH2v2h and the land cover project of the Climate Change Initiative (CCI) program of the European Space 1593 Agency (ESA). The latter is based on purely remotely sensed methods, while the former makes use of national harvest 1594 data from the U.N. Food and Agricultural Organization. 1595 LUH2v2-ESACCI: "We describe here the input data and algorithms used to create the land cover maps specific for 1596 our CMIP6 simulations using the historical/future reconstruction of land use states provided as reference datasets for

CMIP6 within the land use harmonization database LUH2v2h (Hurtt et al., 2020). More details are provided on the devoted web page <u>https://orchidas.lsce.ipsl.fr/dev/lccci</u> which shows further tabular, graphical and statistical data. The overall approach relies on the combination of the LUH2v2 data with present-day land cover distribution derived from satellite observations for the past decades. The main task consists in allocating the land-use types from LUH2v2 in the different PFTs for the historical period and the future scenarios. The natural vegetation in each grid cell is defined as the PFT distribution derived from the ESA-CCI land cover product for the year 2010 to which pasture fraction and

1603 crop fraction from LUH2v2 (for the year 2010) have been subtracted from grass and crop PFTs. This characterization

of the natural vegetation in terms of PFT distribution is assumed invariant in time and is used for both the historicalperiod and the different future scenarios" (Lurton et al., 2020).





1607

Figure A3: A comparison of the version of ORCHIDEE used in previous synthesis of Petrescu et al. (2021b) compared
to the same version using the forcing prepared for this work (ORCHIDEE-V2021) and the version with the coupled
C-N cycle from this work (ORCHIDEE-N-V2021). For the current work, both the version shown with the Europe-





specific nitrogen forcing prepared under VERIFY for the years 1995-2018 (ORCHIDEE-N-V2021) and that using the
standard nitrogen forcing from the N₂O Model Intercomparison Project (NMIP; Tian et al., 2018) as supplied to the
TRENDY model intercomparison is shown (ORCHIDEE-N-V2021 NMIP).

1614

1615 Uncertainty: In the ORCHIDEE model, uncertainty arises from three primary sources: parameters, forcing data 1616 (including spatial and temporal resolution), and model structure. Some researchers argue that the initial state of the 1617 model (i.e., the values of the various carbon and water pools at the beginning of the production run, following model 1618 spinup) represents a fourth area. However, the initial state of this version of ORCHIDEE is defined by its equilibrium 1619 state, and therefore a strong function of the parameters, forcing data, and model structure, with the only independent 1620 choice being the target year of the initial state. Out of the three primary areas of uncertainty, the climate forcing data 1621 is dictated by the VERIFY project itself, thus removing that source from explaining observed differences among the 1622 models, although it can still contribute to uncertainty between the ORCHIDEE results and the national inventories. 1623 The land use/land cover maps, another major source of uncertainty for ORCHIDEE carbon fluxes, have also been 1624 harmonized to a large extent between the bottom-up carbon budget models in the project. Parameter uncertainty and 1625 model structure thus represent the two largest sources of potential disagreement between ORCHIDEE and the other 1626 bottom-up carbon budget models. Computational cost prevents a full characterization of uncertainty due to parameter 1627 selection in ORCHIDEE (and dynamic global vegetation models in general), and uncertainties in model structure 1628 require the use of multiple models of the same type but including different physical processes. Such a comparison has 1629 not been done in the context of VERIFY, although the results from the TRENDY suite of models shown in Fig. 8 give 1630 a good indication of this. Figure A3 shows a small influence from the nitrogen forcing, likely because the European 1631 nitrogen forcing is only available from 1995-2018 and ORCHIDEE carries out almost 500 years of simulation prior 1632 to this point. Many major carbon pools (i.e., woody biomass, soil carbon) have built up a large amount of inertia over 1633 that time and are unlikely to undergo dramatic changes for any realistic forcing over the past. A similar conclusion can be reached from simulations ORCHIDEE-V2019 and ORCHIDEE-V2021 in Fig. A3, which only differ in 1634 1635 meteorological forcing from 1981-2020.

1636

1637 *CABLE-POP*

1638 CABLE-POP (Haverd et al., 2018) is a global terrestrial biosphere model developed around a 1639 core biogeophysics module (Wang & Leuning, 1998) and a biogeochemistry module including cycles of nitrogen and 1640 phosphorus (Wang et al., 2010). Only nitrogen cycling was turned on for the present simulations. The model also 1641 includes modules simulating woody demography (Haverd et al., 2013) as well as land use change and land 1642 management (Haverd et al., 2018). The model distinguishes seven plant functional types which can co-occur in a 1643 given grid cell. CABLE-POP does not simulate (natural) dynamic vegetation and the distribution and cover fraction 1644 of PFTs is only affected by land use change. Forest demography (establishment, age class distribution, mortality) is 1645 accounted for in the simulations, as are natural disturbances and forest management (wood harvest).

1646 For the simulations described here, a baseline land cover map was created from the HILDA+ dataset for the year 19011647 and vegetation classes in the dataset were reclassified to correspond to PFTs represented in CABLE-POP. Land use





transitions as well as land management (harvest) were prescribed from the LUH2v2h dataset over the entire simulation
period. Crops and pastures are treated as C3 grasses but are subject to agricultural harvest fluxes as given by LUH2v2h.
The use of HILDA+ data for the land cover distribution and the LUH2v2h for the representation of land cover/land
use change likely introduced additional uncertainties resulting from a potential mismatch between the two data sets.

1652

1653

1654 CO₂ Emissions from inland waters

In this study we did not update these estimates and they are therefore identical to those in Petrescu et al. 1655 1656 (2021b). These estimates represent a climatology of average annual CO₂ emissions from rivers, lakes and reservoirs 1657 at the spatial resolution of 0.1°. The approach combines CO₂ evasion fluxes from the global river network, as estimated by the empirical model of Lauerwald et al. (2015), with the lakes and reservoirs estimates by Hastie et al. (2019) for 1658 1659 the boreal biome and by Raymond et al. (2013) for the lower latitudes. The Lauerwald et al. and Hastie et al. studies 1660 follow the same approach and rely on the development of a statistical prediction model for inland water pCO_2 at 0.5° 1661 using global, high-resolution geodata. The pCO2 climatology was then combined with different estimates of the gas 1662 transfer velocity k to produce the resulting map of CO₂ evasion. The Raymond et al. study only provides mean flux 1663 densities at the much coarser spatial resolution of the so-called COSCAT regions. All estimates were then downscaled 1664 to 0.1° using the spatial distribution of European inland water bodies. Note that in contrast to Hastie et al. (2019), the 1665 areal distribution of lakes was extracted from the HYDROLAKES database (Messager et al., 2016), to be consistent 1666 with the estimates of inland water N₂O and CH₄ presented by Petrescu et al. (2021b).

1667 Uncertainty: Monte Carlo simulations were performed to constrain uncertainties resulting from both the pCO₂1668 prediction equation and the choice of the k formulation.

1669

1670 *CBM*

1671 The Carbon Budget Model developed by the Canadian Forest Service (CBM-CFS3), can simulate the 1672 historical and future stand- and landscape-level C dynamics under different scenarios of harvest and natural 1673 disturbances (fires, storms), according to the standards described by the IPCC (Kurz et al., 2009). Since 2009, the 1674 CBM has been tested and validated by the Joint Research Centre of the European Commission (JRC), and adapted to 1675 the European forests. It is currently applied to 26 EU Member States, both at country and NUTS2 level (Pilli et al., 1676 2016).

1677Based on the model framework, each stand is described by area, age and land use classes and up to 101678classifiers based on administrative and ecological information and on silvicultural parameters (such as forest1679composition and management strategy). A set of yield tables define the merchantable volume production for each1680species while species-specific allometric equations convert merchantable volume production into aboveground1681biomass at stand-level. At the end of each year the model provides data on the net primary production (NPP), carbon1682stocks and fluxes, as the annual C transfers between pools and to the forest product sector.





1683	The model can support policy anticipation, formulation and evaluation under the LULUCF sector, and it is
1684	used to estimate the current and future forest C dynamics, both as a verification tool (i.e., to compare the results with
1685	the estimates provided by other models) and to support the EU legislation on the LULUCF sector (Grassi et al., 2018a).
1686	In the biomass sector, the CBM can be used in combination with other models, to estimate the maximum wood
1687	potential and the forest C dynamic under different assumptions of harvest and land use change (Jonsson et al., 2018).
1688	Uncertainty: Quantifying the overall uncertainty of CBM estimates is challenging because of the complexity of each
1689	parameter. The uncertainty in CBM arises from three primary sources: parameters, forcing data (including spatial and
1690	temporal resolution) and model structure. It is linked to both activity data and emission factors (area, biomass volume
1691	implied by species specific equation to convert the merchantable volume to total aboveground biomass (used as a
1692	biomass expansion factor)) as well to the capacity of each model to represent the original values, in this case estimated
1693	through the mean percentage difference between the predicted and observed values. A detailed description of the
1694	uncertainty methodology is found in Pilli et al. (2017).
1695	

1696 Explanatory note on the extrapolation of Net Biome Productivity for the period 2017-2020 (Matteo Vizzarri,
1697 Roberto Pilli, Giacomo Grassi, EC-JRC)

- 1698 Background
- 1699 We performed a linear extrapolation of forest Net Biome Productivity (NBP) by country (EU 25 Member States and
- 1700 UK) in the period 2017-2020 based on the correlation between NBP and harvest from the period 2000-2015. Cyprus
- and Malta are excluded from the analysis because of missing historical data.
- 1702 Input data
- 1703 Table A4 reports a summary of input data sources.
- 1704 *Table A4: main input data used in the extrapolation of NBP for the period 2017-2020.*

	Uni t	Temporal resolution	Source
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Wood removals (HWP pool)	t C	Annual (2000-2015)	CBM calibration run
Forest area	ha	Annual (2000-2020)	FAOSTAT ²³
Roundwood amount	m ³	Annual (2000-2020)	FAOSTAT ²⁴
NBP	t C	Annual (2000-2015)	CBM calibration run

1705

1706 Assessment procedure

1707 The extrapolation of the NBP for the period 2017-2020 was obtained throughout the following steps:

1708	1.	For each country (EU 25 Member States + UK), we first calculated the average conversion factor -
1709		representing a correspondence between one ton of biomass carbon removed and one cubic meter of wood per
1710		hectare – for the period 2000-2015 through equation [1]:

1711
$$CF_{2000-2015} = \sum_{t=2000}^{2015} \frac{HWP_t}{\frac{RW_t}{A_{2015}}} \quad \text{eq. (1)}$$

1712where: $CF_{2000-2015}$ is the average conversion factor per hectare in the period 2000-2015 (t C m⁻³ ha1713¹); HWP_t is the carbon content per ha in harvested wood products in year t (t C year⁻¹), as derived from the CBM1714model run; RW is the total roundwood removals in year t (m³ year⁻¹) (source: FAOSTAT²⁵); A_{2015} is the managed1715forest area in year 2015 (ha; source: Forest Europe 2015).

1716
2. Using the average conversion factor estimated in eq. 1, we converted, for each country, the total roundwood removals per ha derived from FAOSTAT for the period 2017 2020, to the corresponding amount of carbon removals per ha, through equation [2]:

1719 1720

$$HWP_{conv} = CF_{2000-2015} \cdot \left(\frac{RW_t}{A_{2015}}\right) \left[\forall t = 2017 \div 2020\right] \qquad \text{eq. (2)}$$

1721where: HWP_{conv} is the amount of carbon removals per hectare in year t (t C ha⁻¹ year⁻¹); $CF_{2000-2015}$ is the1722average conversion factor per hectare in the period 2000-2015 (t C m⁻³ ha⁻¹); RW_t is the total roundwood in year t1723(m³ year⁻¹) (source: FAOSTAT²⁶); A_{2015} is the managed forest area in the year 2015 (ha).

1724	3.	Then, for each country and the period 2000-2015, we performed a linear regression to search for significant
1725		correlation between the harvest amount (i.e. HWP in t C ha ⁻¹ yr ⁻¹) and NBP, according to the generalized
1726		equation:
1727		$NBP = a + b \cdot (HWP)$ eq. (3)

In this case, we assumed NBP as the dependent variable (t C ha⁻¹ year⁻¹), the amount of harvest (t C ha⁻¹ year⁻¹)
as the main driver affecting the short term evolution of NBP, in absence of other exogenous natural disturbances; *a* is the intercept of the linear trendline; *b* is the coefficient of the independent variable harvest

²³ https://www.fao.org/faostat/en/#data/RL

²⁴ <u>https://www.fao.org/faostat/en/#data/FO</u>

²⁵ https://www.fao.org/faostat/en/#data/FO

²⁶ https://www.fao.org/faostat/en/#data/FO





- 1731amount (i.e. HWP) (m³ ha⁻¹ year⁻¹). This approach is consistent with the methodological assumptions reported1732in Jonsson et al. (2021).
- 4. We finally calculated the **NBP in the period 2017-2020** for each country through equation [4]:

1734
$$NBP_{t,m} = (a + b \cdot HWP_{conv})_{t,m} \quad \text{eq. (4)}$$

- 1735 where: $NBP_{t,m}$ is the Net Biome Productivity for year t, country m (t C ha⁻¹ year⁻¹); $a_{t,m}$ is the intercept 1736 of the linear trendline for year t, country m; $b_{t,m}$ is the coefficient of the independent variable in the trendline; 1737 $HWP_{conv(t,m)}$ is the amount of carbon removal per ha for year t, country m (t C ha⁻¹ year⁻¹).
- 1738 Forest area and parameters used in equation [4] by country are reported in Table A5.
- 1739Table A5: country-based forest area in 2015 and parameters used in equation [4]. *: significant (p<0.05); NS: not</th>1740significant (p>0.05).

EU 25 + UK	CF (2000- 2015)	Intercept (a)	Coefficient (b)	<i>p</i> <0.0 5
Austria	0.28	2.60	-1.57	*
Belgium	0.18	2.97	-1.54	*
Bulgaria	0.22	1.17	-2.13	*
Croatia	0.28	1.42	-1.27	*
Czechia	0.22	2.55	-1.21	*
Denmark	0.16	1.92	-1.21	*
Estonia	0.20	1.16	-1.08	*
Finland	0.23	1.15	-1.20	*





EU 25 + UK	CF (2000- 2015)	Intercept (a)	Coefficient (b)	<i>p</i> <0.0 5
France	0.19	1.63	-1.17	*
Germany	0.21	2.55	-1.23	*
Greece	0.20	1.17	-1.75	ns
Hungary	0.27	1.50	-1.54	*
Ireland	0.18	6.12	-5.45	*
Italy	0.23	0.69	0.39	ns
Latvia	0.19	2.00	-1.77	*
Lithuania	0.22	1.11	-0.89	*
Luxembourg	0.20	1.79	-1.40	*
Netherlands	0.22	2.44	-2.01	*
Poland	0.21	2.49	-2.16	*
Portugal	0.29	1.39	-1.01	*
Romania	0.32	1.54	-1.65	*
Slovakia	0.28	2.57	-1.42	*
Slovenia	0.24	2.07	-1.55	*
Spain	0.28	0.26	0.18	ns
Sweden	0.23	1.02	-1.20	*


EU 25 + UK	EU 25 + UK CF (2000-		Coefficient	<i>p</i> <0.0
	2015)		(b)	5
United Kingdom	0.19	2.27	-1.34	*

1741

1742 *Additional notes*

1743 Because of biased estimates, values for the year 2016 were excluded from this analysis.

1744 Extrapolated NBP for Czech Republic, Ireland and Netherlands were negative (thus showing emissions) because of

an increase of harvest in the corresponding years (2017-2020) compared to the previous period 2000-2015. Estonia

shows negative extrapolated NBP only for the year 2018.

1747 EFISCEN-Space

1748 The European Forest Information SCENario Model (EFISCEN) is a large-scale forest model that projects 1749 forest resource development on a regional to European scale. The model uses aggregated national forest inventory 1750 data as a main source of input to describe the current structure and composition of European forest resources. The 1751 model projects the development of forest resources, based on scenarios for policy, management strategies and climate 1752 change impacts. With the help of biomass expansion factors, stem wood volume is converted into whole-tree biomass 1753 and subsequently to whole tree carbon stocks. Information on litter fall rates, felling residues and natural mortality is 1754 used as input into the soil module YASSO (Liski et al., 2005), which is dynamically linked to EFISCEN and delivers 1755 information on forest soil carbon stocks. The core of the EFISCEN model was developed by Prof. Ola Sallnäs at the 1756 Swedish Agricultural University (Sallnäs, 1990). It has been applied to European countries in many studies since then, 1757 dealing with a diversity of forest resource and policy aspects. A detailed model description is given by Verkerk et al. 1758 (2016), with online information on availability and documentation of EFISCEN at http://efiscen.efi.int. The model 1759 and its source code are freely available, distributed under the GNU General Public License conditions 1760 (www.gnu.org/licenses/gpl-3.0.html).

1761 In this report the follow-up of the EFISCEN model was used, called EFISCEN-Space (Schelhaas et al., in 1762 prep). EFISCEN-Space simulates the development of the forest at the level of the plots as measured in the national 1763 forest inventories, thereby providing a much higher spatial detail. The simulation is based on the distribution of trees 1764 over diameter classes rather than age as in the old EFISCEN model. This allows the simulation of a wider variety of 1765 stand structures, species mixtures and management options. Similar to the EFISCEN model, biomass expansion 1766 factors and the YASSO soil carbon model are used to provide carbon balances for the forest. For use within VERIFY, 1767 individual plot results are aggregated to a 0.125 degree grid. For the moment only 15 European member states are 1768 included, partly due to the lack of an appropriate national forest inventory in the other member states, or because the 1769 data could not be shared. No formal sensitivity and uncertainty analysis has been conducted yet.





1770 Figure 5 shows results which vary from year-to-year. In practice, the model was initialized with starting 1771 years depending on the country, assuming that all data applied to this year. The model then produced stock and flux 1772 changes for the subsequent five-year period, reporting a single mean value per pixel. To compute timeseries for the 1773 EU27+UK, it was further assumed that these values were valid across 2005-2020. As the fluxes were given per square 1774 meter of forest, they were scaled by the total area of the forest in each pixel found on the land use/land cover maps 1775 used by the ORCHIDEE DGVM. This explains why the numbers vary from year to year; the flux per square meter of 1776 forest does not change, but the total amount of forest area changes slightly. It should be noted that country-level values 1777 available on the VERIFY website are only available for the five-year period for which the model produces a mean 1778 result.

1779

Uncertainties: A sensitivity analysis of EFISCEN v3 is described in detail in Chapter 6 of the user manual (Schelhaas
et al., 2007). Total sensitivity is caused by especially young forest growth, width of volume classes, age of felling and
few other variables. Scenario uncertainty comes on top of this when projecting in future. Within VERIFY, a full
uncertainty analysis has been completed, enabling the estimation of uncertainty ranges of the various output variables
(Schelhaas et al., 2020).

1785

1786 EPIC-IIASA

1787 The Environmental Policy Integrated Climate (EPIC) model is a field-scale process-based model (Izaurralde 1788 et al., 2006; Williams, 1990) which calculates, with a daily time step, crop growth and yield, hydrological, nutrient 1789 and carbon cycling, soil temperature and moisture, soil erosion, tillage, and plant environment control. Potential crop 1790 biomass is calculated from photosynthetically active radiation using the radiation-use-efficiency concept modified for 1791 vapor pressure deficit and atmospheric CO₂ concentration effect. Potential biomass is adjusted to actual biomass 1792 through daily stress caused by extreme temperatures, water and nutrient deficiency, or inadequate aeration. The 1793 coupled organic C and N module in EPIC (Izaurralde et al., 2006) distributes organic C and N between three pools of 1794 soil organic matter (active, slow and passive) and two litter compartments (metabolic and structural). EPIC calculates 1795 potential transformations of the five compartments as regulated by soil moisture, temperature, oxygen, tillage and 1796 lignin content. Daily potential transformations are adjusted to actual transformations when the combined N demand 1797 in all receiving compartments exceeds the N supply from the soil. The transformed components are partitioned into 1798 CO2 (heterotrophic respiration), dissolved C in leaching (DOC) and the receiving SOC pools. EPIC also calculates 1799 SOC loss with erosion.

1800The EPIC-IIASA (version EU) modeling platform was built by coupling the field-scale EPIC version 08101801with large-scale data on land cover (cropland and grasslands), soils, topography, field size, crop management practices1802and grassland cutting intensity aggregated at a 1x1 km grid covering European countries (Balkovič et al., 2018, 2013).1803In VERIFY, a total of 10 major European crops including winter wheat, winter rye, spring barley, grain maize, winter1804rapeseed, sunflower, sugar beet, potatoes, soybean and rice were used to represent agricultural production systems in1805European cropland. Crop fertilization and irrigation were estimated for NUTS2 statistical regions between 1995 and18062010 (Balkovič et al., 2013). For VERIFY, the simulations were carried out assuming conventional tillage, consisting





1807 of two cultivation operations and moldboard plowing prior to sowing and an offset disking after harvesting of cereals. 1808 Two row cultivations during the growing season were simulated for maize and one ridging operation for potatoes. It 1809 was assumed that 20 % of crop residues are removed in the case of cereals (excluding maize), while no residues are 1810 harvested for other crops.

1811 A total of five managed grassland types with distinct temperature requirements, biomass productivity, and 1812 phenology were used to represent the C-cycle in European grasslands. High-productive generic winter pasture and tall 1813 fescue-based grasslands were used for Atlantic Europe, low fescue grasslands for the cool climates of Nordic regions 1814 and high mountains, high-productive tall fescue-based grasslands and low-productive bluegrass types for continental 1815 Europe, and low-productive bromegrass and high-productive winter pastures in the Mediterranean regions. Annual 1816 nitrogen and carbon inputs, including inorganic and manure fertilization, and atmospheric N deposition, were obtained 1817 from ISIMIP3 (Jägermeyr et al., 2021). In this dataset, the annual manure production and the fraction of manure from 1818 livestock applied to cropland and rangeland were used from Zhang et al. (2017). The original manure data were re-1819 gridded to half-degree spatial resolution in ISMIP3. In the model, manure is applied as an organic fertilizer with a 1820 C:N ration of 14.5:1. The organic carbon and nitrogen are added to the fresh organic litter pool where they decompose 1821 in a manner identical to the fresh litter from vegetation, while mineral N from manure is added to the soil nitrate and 1822 ammonium pools. The distribution of herbage biomass export intensity was constructed based on (Chang et al., 2016). 1823 Uncertainty: In EPIC, uncertainties arise from three primary sources which were described in detail by ORCHIDEE. 1824 A detailed sensitivity and uncertainty analysis of EPIC-IIASA regional carbon modeling is presented in (Balkovič et 1825 al., 2020).

1826

1827 ECOSSE (grasslands)

1828 ECOSSE is a biogeochemical model that is based on the carbon model ROTH-C (Jenkinson and Rayner, 1829 1977; Jenkinson et al. 1987; Coleman and Jenkinson, 1996) and the nitrogen-model SUNDIAL (Bradbury et al., 1993; 1830 Smith et al., 1996). All major processes of the carbon and nitrogen dynamics are considered (Smith et al., 2010a,b). 1831 Additionally, in ECOSSE processes of minor relevance for mineral arable soils are implemented as well (e.g., methane 1832 emissions) to have a better representation of processes that are relevant for other soils (e.g., organic soils). ECOSSE 1833 can run in different modes and for different time steps. The two main modes are site specific and limited data. In the 1834 later version, basic assumptions/estimates for parameters can be provided by the model. This increases the uncertainty 1835 but makes ECOSSE a universal tool that can be applied for large scale simulations even if the data availability is 1836 limited. To increase the accuracy in the site-specific version of the model, detailed information about soil properties, 1837 plant input, nutrient application and management can be added as available.

1838During the decomposition process, material is exchanged between the SOM pools according to first order1839rate equations, characterized by a specific rate constant for each pool, and modified according to rate modifiers1840dependent on the temperature, moisture, crop cover and pH of the soil. The model includes five pools with one of1841them being inert. The N content of the soil follows the decomposition of the SOM, with a stable C:N ratio defined for1842each pool at a given pH, and N being either mineralized or immobilized to maintain that ratio. Nitrogen released from1843decomposing SOM as ammonium (NH4+) or added to the soil may be nitrified to nitrate (NO3-).





For spatial simulations the model is implemented in a spatial model platform. This allows to aggregate the input parameter for the desired resolution. ECOSSE is a one-dimensional model and the model platform provides the input data in a spatial distribution and aggregates the model outputs for further analysis. While climate data are interpolated, soil data are represented by the dominant soil type or by the proportional representation of the different soil types in the spatial simulation unit (this is in VERIFY a grid cell).

1849 Uncertainty: In ECOSSE, uncertainty arises from three primary sources: parameters, forcing data (including spatial1850 and temporal resolution), and model structure. These uncertainties are not yet quantified.

1851

1852 Bookkeeping models

1853 We make use of data from two bookkeeping models: BLUE (Hansis et al., 2015) and H&N (Houghton &
1854 Nassikas, 2017).

1855 The BLUE model provides a data-driven estimate of the net land use change fluxes. BLUE stands for 1856 "bookkeeping of land use emissions". Bookkeeping models (Hansis, 2015; Houghton, 1983) calculate land-use change 1857 CO2 emissions (sources and sinks) for transitions between various natural vegetation types and agricultural lands. The 1858 bookkeeping approaches keep track of the carbon stored in vegetation, soils, and products before and after the land-1859 use change. In BLUE, land-use forcing is taken from the Land Use Harmonization, LUH2, for estimates within the 1860 annual global carbon budget. The model provides data at annual time steps and 0.25 degree resolution. Temporal 1861 evolution of carbon gain or loss, i.e., how fast carbon pools decay or regrow following a land-use change, is based on 1862 response curves derived from literature. The response curves describe decay of vegetation and soil carbon, including 1863 transfer to product pools of different lifetimes, as well as carbon uptake due to regrowth of vegetation and subsequent 1864 refilling of soil carbon pools. In this report we present two versions of BLUE: BLUEvVERIFY and BLUEvGCP. The 1865 BLUEvVERIFY version is a set of runs made for VERIFY, using the Hilda+²⁷ product (Ganzenmüller et al., 2022).

1866 The H&N model (Houghton et al., 1983) calculates land-use change CO₂ emissions and uptake fluxes for 1867 transitions between various natural vegetation types and agricultural lands (croplands and pastures). The original 1868 bookkeeping approach of Houghton (2003) keeps track of the carbon stored in vegetation and soils before and after 1869 the land-use change. Carbon gain or loss is based on response curves derived from literature. The response curves 1870 describe decay of vegetation and soil carbon, including transfer to product pools of different life-times, as well as 1871 carbon uptake due to regrowth of vegetation and consequent re-filling of soil carbon pools. Natural vegetation can 1872 generally be distinguished into primary and secondary land. For forests, a primary forest that is cleared can never 1873 return back to its original carbon density. Instead, long- term degradation of primary forest is assumed and represented 1874 by lowered standing vegetation and soil carbon stocks in the secondary forests. Apart from land use transitions 1875 between different types of vegetation cover, forest management practices in the form of wood harvest volumes are 1876 included. Different from dynamic global vegetation models, bookkeeping models ignore changes in environmental 1877 conditions (climate, atmospheric CO₂, nitrogen deposition and other environmental factors). Carbon densities at a 1878 given point in time are only influenced by the land use history, but not by the preceding changes in the environmental

²⁷https://landchangestories.org/hildaplus/





state. Carbon densities are taken from observations in the literature and thus reflect environmental conditions of thelast decades. In this study an updated H&N version submitted to the GCP2021 is used.

1881 Uncertainty: Uncertainties can be captured through simulations varying uncertain parameters, input data, or process

representation. A large contribution of uncertainty can be expected from various input datasets. Apparent uncertainties
arise from the land-use forcing data (Gasser et al., 2020; Hartung et al., 2021; Ganzenmüller et al., 2022), the
equilibrium carbon densities of soil and vegetation and allocation of material upon a land-use transition (Bastos et al.,

1885 2021), and the response curves built to reflect carbon pool decay and regrowth after land-use transitions. Furthermore,

1886 studies have shown that different accounting schemes (Hansis et al., 2015) and initialization settings at the start of the

1887 simulations (Hartung et al., 2021) lead to different emission estimates even decades later.

1888

1889 *FAOSTAT*

1890 FAOSTAT: Statistics Division of the Food and Agricultural Organization of the United Nations provides 1891 updates for the LULUCF CO_2 emissions for the period 1990-2019, available at: 1892 https://www.fao.org/faostat/en/#data/GT and its sub-domains. The FAOSTAT emissions land use database is 1893 computed following a Tier 1 approach of IPCC (2006). Geospatial data are the source of AD for the estimates of 1894 emissions from cultivation of organic soils, biomass and peat fires. GHG emissions are provided by countries, regions 1895 and special groups, with global coverage, relative to the period 1990-present (with annual updates). Land Use Total 1896 contains all GHG emissions and removals produced in the different Land Use sub-domains, representing four IPCC 1897 Land Use categories, of which three land use categories: forest land, cropland, grassland and biomass burning. LULUCF emissions consist of CO2 associated with land use and change, including management activities. CO2 1898 1899 emissions/removals are computed at Tier 3 using carbon stock change. To this end, FAOSTAT uses Forest area and 1900 carbon stock data from FRA (2015), gap-filled and interpolated to generate annual time-series. As a result CO₂ 1901 emissions/removals are computed for forest land and net forest conversion, representing respectively IPCC categories 1902 "Forest land" and "Forest land converted to other land uses". CO2 emissions are provided as by country, regions and 1903 special groups, with global coverage, relative to the period 1990-most recent available year (with annual updates), 1904 expressed as net emissions/removals as Gg CO2, by underlying land use emission sub-domain and by aggregate (land 1905 use total).

1906 Uncertainty: FAOSTAT uncertainties are not available.

1907

1908 TRENDY DGVMs

1909 The TRENDY (Trends in net land-atmosphere carbon exchange over the period 1980-2010) project 1910 represents a consortium of dynamic global vegetation models (DGVMs) following identical simulation protocols to 1911 investigate spatial trends in carbon fluxes across the globe over the past century. As DGVMs, the models require 1912 climate, carbon dioxide, and land use change input data to produce results. In TRENDY, all three of these are 1913 harmonized to make the results across the whole suite of models more comparable. In the case of VERIFY, 15 of the 1914 16 models for TRENDY v10 (except for ISAM) were used. While describing the details of all the models used here





1915	is clearly not possible, DGVMs calculate prognostic variables (i.e., a multitude of C, H ₂ O and energy fluxes) from the
1916	following environmental drivers: air temperature, wind speed, solar radiation, air humidity, precipitation and
1917	atmospheric CO2 concentration. As the run progresses, vegetation grows on each pixel, divided into generic types
1918	which depend on the model (e.g., broadleaf temperate forests, C3 crops), which cycle carbon between the soil, land
1919	surface, and atmosphere, through such processes such as photosynthesis, litter fall, and decay. Limited human
1920	activities are included depending on the model, typically removing aboveground biomass on an annual basis.
1921	Among other environmental indicators, DGVMs simulate positive and negative CO2 emissions from plant
1922	uptake, soil decomposition, and harvests across forests, grasslands, and croplands. Activity data is based on land use
1923	and land cover maps and generally follows Approach 1 as described by the IPCC 2006 guidelines (enabling calculation
1924	of only net changes from year to year). For TRENDY, pixel land cover/land use fractions were based on the land use
1925	map LUH2 (Hurtt et al., 2020) and the HYDE land-use change data set (Klein Goldewijk et al., 2017a, b). Both of
1926	these maps rely on FAO statistics on agricultural land area and national harvest data.
1927	Uncertainty: In TRENDY v10 uncertainties are model specific and described by Friedlingstein et al. (2022). The
1928	spread of the 15 TRENDY models used by this study (Fig. 8) gives an idea of the uncertainty due to model structure
1929	in dynamic global vegetation models, as the forcing data was harmonized for all models.
1930	
1931	Emissions from lateral transport of carbon (crops, wood, and inland waters)
1932	
1933	Production and consumption of carbon do not always occur on the same grid points. This is particularly relevant for
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1933 1934 1935 1936 1937 1938 1939 1940 1941 1942 1943	Production and consumption of carbon do not always occur on the same grid points. This is particularly relevant for the land surface in the case of crops, wood products, and carbon transfers through the inland water network. The purpose of the work here is primarily to convert the flux changes of the top-down inversions into NGHGI-like stock changes. To convert the flux changes of the inversions (where a positive number represents a flux to the atmosphere, i.e., a source) into NGHGI-like stock changes, one needs to add the crop sink and remove the crop source. The crop sink comes from production numbers in the FAO food balance sheets, while the source is estimated by production plus import minus export (all from the FAO food balance sheets), and both terms make use of conversion factors for each commodity. We take the forestry balance sheets of FAO (production, import and export per commodity), and convert to C mass. For a given year, the fraction of this mass that is released later in the atmosphere in each country is modeled with an e-folding decrease driven by experimental data per country (Mason Earles et al., 2012). Lateral transfers of carbon through inland waters also need to be removed from the inversion results as the terrestrial
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1951 Carbon fluxes to the atmosphere from rivers and lakes were obtained from maps described in Zscheischler 1952 et al. (2017). These methods are similar to those described previously in Petrescu et al. (2021b). The primary difference 1953 is that the updated estimates include smaller lakes and reservoirs not represented in the Global Lakes and Wetland 1954 Database through the use of a scaling law, in addition to the older results being created specifically for Europe, while 1955 the newer results are part of a global product. The emissions from the previous work totaled 25.5 Tg C yr-1 for the EU27+UK, while those used here are 19.8 Tg C yr-1 (with no variability from year-to-year). This difference is 1956 1957 therefore small compared to the river C export, which is included this year for the first time and averages -73.8 Tg for 1958 the period 1990-2020.

1959 One important difference between the fluvial carbon exports reported here and those from a previous work 1960 (Ciais et al., 2021) are that those reported here are rescaled to reasonable global flux reflecting bias in inter-1961 hemispheric exchange. Similar to Bastos et al. (2020), the dissolved organic carbon (DOC) and particulate organic 1962 carbon (POC) exports were rescaled per basin to match the estimates of Resplandy et al. (2018). The global total 1963 organic C was finally rescaled to 500 Tg C/yr, which is considered a reasonable global number based on different 1964 reviews and synthesis efforts (Regnier et al., 2013).

- 1965
- 1966

1967 Top-down CO₂ emissions estimates

1968 CarboScope-Regional

1969 CarboScopeRegional (CSR) (Munassar et al., 2022): CSR is a Bayesian Framework inversion system that employs 1970 a-priori knowledge of the surface-atmosphere carbon fluxes to regularize the solution of the ill-posed inverse problem 1971 arising from the sparseness of observations sampled over limited geographical locations throughout the domain of 1972 interest. Due to the heterogeneity of biogenic fluxes, the convention in CSR is to optimize Net Ecosystem Exchange 1973 (NEE) against measurements of CO₂ dry model fraction at 3-hourly temporal and 0.5° horizontal resolutions, while 1974 ocean fluxes and anthropogenic emissions are prescribed given their better knowledge available compared with NEE. 1975 The prior flux uncertainty is assumed to have a uniform shape in space and time and its spatial correlation is fitted to 1976 a hyperbolic decay function following the assumption of Kountouris et al. (2018a, b). Model-data mismatch 1977 uncertainty is defined weekly in the measurement covariance matrix varying over sites from 0.5 to 4 (ppm) according 1978 to the ability for atmospheric transport models to sample the true concentration at such locations (Rödenbeck, 2005). This uncertainty implicitly encompasses the combinations of atmospheric transport, representation, and measurement 1979 1980 errors and is assumed to be independent at different locations. To separate the lateral influences originating from 1981 outside of the regional domain, the two-step scheme inversion (Rödenbeck et al., 2009) is applied to run a global 1982 inversion with the Eulerian model TM3 at coarse resolutions to provide the lateral boundary conditions to the regional 1983 inversion. In the regional inversion runs, the Lagrangian model STILT (Lin et al., 2003), forced by IFS data from 1984 ECMWF, is used to calculate the surface sensitivities "footprints" over the regional site network (receptors) at hourly 1985 temporal and 0.25° spatial resolutions. Typically, the prior fluxes of CO₂ are obtained from bottom-up model 1986 estimations. Thus, the diagnostic biosphere model VPRM calculates the biogenic fluxes at hourly temporal resolution





preserving the diurnal cycle. Ocean fluxes are obtained from the CarboScope ocean-based fluxes developed in-house
by Rödenbeck et al. (2014). Emissions of fossil fuel are taken from EDGAR_v4.3 inventories updated every year
based on the British Petroleum statistics (BP), and are distributed in space and time using the COFFEE approach
(Steinbach et al., 2011) according to fuel-type and sector.

1991 The v2021 CSR inversions underwent updates in comparison with the previous v2019:

- v2019 from Petrescu et al. (2021b) excluded observations from two sites: La Muela (LMU) in Spain because of inconsistent datasets between releases, and Finokalia (FKL) in Greece due to errors in the dataset. These exclusions resulted in a larger C sink from 2013 onwards (Fig. 9, upper plot). FKL observations start at this time and are the dominant impact over south-east Europe, as it is the only site located there. In v2021 inversions, we included corrected datasets from the FKL site.
- Two new flask sites were included in the v2021 inversions: Shetland Islands in the UK and Centro
 Investigacion Baja in Spain. These sites are also used in the CarboScope global inversion that provides the
 far-field contributions to the EU domain.

2000

2001 Uncertainty: Uncertainties from top-down (TD) estimates can be reported as posterior Bayesian uncertainties. Following the methodology of Chevallier et al. (2007) the CSR inversion system computed maps of uncertainty 2002 2003 reductions for 2006 and 2018 (Fig. A4). The reduction is carried out through an ensemble of 40 members of inversions 2004 using error realizations following a Monte Carlo (MC) approach. Circles on maps refer to locations of stations. In the 2005 inversion system, a MC method is used to generate N ensembles of realizations of prior errors and model-data 2006 mismatch errors. The inversion is repeated for each ensemble member starting from each set of prior and model-data 2007 mismatch errors to generate posterior fluxes. The posterior uncertainty is calculated as the spread over the optimized 2008 fluxes across the whole ensemble. The uncertainty reduction is then calculated as 1- (opost / oprior). It is clear that 2009 larger ensembles will lead to better convergence of the error reduction. However, due to computational limitations, 40 2010 ensemble members were selected as a good compromise.









2013 Figure A4: CSR uncertainty reduction maps computed as $1 - (\sigma_{post} / \sigma_{prior})$ for 2006 and 2018 using a Monte Carlo 2014 approach focused on prior errors. The circles represent the observation stations network.

2015 Figure A4 represents a preliminary attempt at how the inclusion of additional observation stations (additional 2016 circles in the right-side figure for Germany, Switzerland, Finland compared to the left-side figure) might reduce the 2017 uncertainty. However, the two different simulation years (2006 and 2018) might also differ in terms of other factors 2018 which may lead to lower uncertainties in a given year (e.g., climatological conditions, such as the 2018 drought year). 2019 Several caveats remain. When comparing the uncertainty over pixels or subregions in the domain of interest, 2020 the maps of uncertainty reduction should be interpreted together with the maps of posterior uncertainty to give a better 2021 illustration of the magnitude of uncertainty. The maps of uncertainty reduction reflect only the random uncertainties. 2022 The systematic uncertainties are still poorly characterized, including uncertainties due to atmospheric transport 2023 modeling, dependence on the prior fluxes, and the weighting between the prior and observation uncertainties. To 2024 improve knowledge of the systematic uncertainties, dedicated studies with controlled comparisons between inversions 2025 using different atmospheric transport models (such as planned with the Community Inversion Framework, Berchet et 2026 al., 2021) are still needed. Furthermore, the posterior uncertainty and uncertainty reductions between inversions 2027 depend on internal parameterizations, e.g., the weighting of prior and observation uncertainties. Future efforts should focus on establishing best practices on how to set-up inversions and quantification of systematic uncertainties, 2028 2029 including as well tests of the fidelity of models against data (Simmonds et al., 2021).

2030

2031

2032 *LUMIA*

2033The LUMIA inversion system (Monteil and Scholze, 2021) is a regional atmospheric inversion system, which2034was designed to produce estimates of the land-atmosphere carbon exchanges based on in-situ CO2 observations from2035the ICOS network. It relies on the FLEXPART 10.4 Lagrangian transport model (Pisso et al., 2019) to compute the2036transport of CO2 fluxes within a regional domain (15°W; 33°N to 35°E, 73°N) at a 0.5°, 3-hourly resolution. Boundary





conditions are provided in the form of timeseries of far-field contributions at the observation sites, obtained from a
global TM5-4DVAR inversion (using the 2-step inversion approach of Rödenbeck et al., 2009). Both transport models
were driven by ECMWF ERA-Interim data, up to 2018, and by ECMWF ERA5 data afterwards.

2040 The inversions solve for weekly offsets to the prior NEE/NBP estimate, at a variable spatial resolution, 2041 highest where the observational coverage is better (up to 0.5° upwind of the observation sites). The optimal solution 2042 is searched for using a variational inversion approach (preconditioned conjugate gradient). The inversions were constrained by in-situ and flask observations from 66 European 2043 2044 observation sites, although only a subset of these sites is usually available at a 2045 given time. The observation uncertainties were set to 1 ppm/week at all sites (the 2046 uncertainty of a single observation is therefore higher, on average 5.2 ppm, and given 2047 by \sqrt{n} , with n the number of assimilated observations at the same site in a ±3.5 day window around the observation time). The prior NEE was produced using the LPJ-2048 2049 GUESS model (Smith et al., 2014), driven by ECMWF ERA5 meteorological data.

The inversion also accounts for (prescribed) anthropogenic CO₂ fluxes from the EDGAR/TNO product (https://doi.org/10.18160/Y9QV-S113) and for atmosphere-ocean CO₂ exchanges from the Jena-CarboScope oc_v2021 product (https://www.bgc-jena.mpg.de/CarboScope/oc/oc_v2021.html). The uncertainties on the prior NEE were set proportional to the sum of the absolute value of the 3-hourly fluxes in each 7-day optimization interval (so the uncertainty is not zero even if the net flux is zero), and scaled to a total value of 0.45 PgC/year, accounting for covariances based on Gaussian (spatial) and exponential (temporal) correlation decay functions, with correlation lengths of respectively 500 km and 1 month (see Monteil and Scholze, 2021, for details).

The main differences from the LUMIA setup used in Thompson et al. (2014) are the specification of prior and observation uncertainties (here made, on purpose, more comparable to those used in the CSR inversions), and the implementation of flux optimization at a variable spatial resolution (which has negligible impact on the results but improves the model performance).

2061 CIF-CHIMERE - land CO₂

2062 CIF-CHIMERE is used for both CO₂ land and CO₂ fossil emission estimates, and this section only describes
 2063 the CO₂ land estimates.

The CIF-CHIMERE inversions have been generated with the variational mode of the Community Inversion Framework (CIF, Berchet et al., 2021) coupled to the regional Eulerian atmospheric chemistry-transport model CHIMERE (Menut et al., 2013; Mailler et al., 2017) and to its adjoint code. They are set-up in a manner that is close to that of the PYVAR-CHIMERE inversions of Broquet et al. (2013), of Thompson et al. (2020) and of Monteil et al. (2020).

A European configuration of CHIMERE is used; this configuration covers latitudes 31.75-73.25°N and
 longitudes 15.25°W -34.75°E with a 0.5°×0.5° horizontal resolution and 17 vertical layers up to 200 hPa.





2071 Meteorological forcing for CHIMERE is generated using the European Center for Medium Range Weather 2072 Forecasting (ECMWF) operational forecasts. Initial, lateral and top boundary conditions for CO₂ concentrations are 2073 generated from the new CAMS global CO₂ inversions v20r2 (Chevallier et al., 2010).

2074 The inversion assimilates in situ CO₂ data from continuous measurements stations compiled in the VERIFY 2075 Deliverable D3.12 and in the Table A1 from the VERIFY CIF Inversion Protocol (Thompson et al., 2021). More 2076 specifically, the inversion assimilates 1-hour averages of the measured CO2 mole fractions during the time window 2077 12:00-18:00 UTC for low altitude stations (below 1000 masl) and 0:00-6:00 UTC for high altitude stations (above 2078 1000 masl). The inversion optimizes 6-hourly mean NEE and ocean fluxes at the 0.5°×0.5° resolution of CHIMERE. 2079 The anthropogenic CO₂ emissions, considered as perfect and consequently not optimized in the inversions, are based 2080 on the spatial distribution of the EDGAR-v4.2 inventory, on national and annual budgets from the BP (British 2081 Petroleum) Statistics and on temporal profiles at hourly resolution derived with the COFFEE approach (Steinbach et 2082 al., 2011).

2083 The prior estimate of NEE and its uncertainty covariance matrix are specified using ORCHIDEE model 2084 simulations of NEE and respiration, respectively, following the general approach of Broquet et al. (2011). The 2085 temporal and spatial correlation scales for the prior uncertainty in NEE are set to ~ 1 month and 200 km (following the 2086 diagnostics of Kountouris et al., 2015), with no correlation between the four 6-hour windows of the same day. The 2087 ocean prior fluxes come from a hybrid product of the University of Bergen coastal ocean flux estimate and the 2088 Rödenbeck global ocean estimate (Rodenbeck et al., 2014). Fluxes from biomass burning are ignored. The observation 2089 error covariance matrix is set-up to be diagonal, ignoring the correlations between errors for different hourly averages 2090 of the CO₂ measurements (which has been justified by the analysis of Broquet et al., 2011). The variances for hourly 2091 data are based on the values from Broquet et al. (2013), which vary depending on the sites and season, and which are 2092 derived from Radon model-data comparisons.

About 12 iterations are needed to reduce the norm of the gradient of J by 95 %, using the M1QN3 limited memory quasi-Newton minimisation algorithm (Gilbert et Lemaréchal, 1989). To cover the whole analysis period (2005-2020), a series of 7-month (including an overlapping of 15 days between consecutive periods) inversions is performed. Posterior estimates of NEE at 1-hourly and $0.5^{\circ} \times 0.5^{\circ}$ spatial resolution are generated for the full period of analysis.

2098

2099 Uncertainty: Estimates of the uncertainty of regional inversions over Europe can be found by comparing against the
 2100 results of the other regional inversions in this work (the ensembles of EUROCOM, CarboScopeRegional, and
 2101 LUMIA).

2102

2103 GCP 2021

Top-down estimates of land biosphere fluxes are provided by a number of different inverse modeling systems that use atmospheric concentration data as input, as well as prior information on fossil emissions, ocean fluxes, and land biosphere fluxes. The land biosphere fluxes, and in some systems the ocean fluxes, are estimated using a statistical optimization involving atmospheric transport models. The inversion systems differ in the transport models used,





optimization methods, spatiotemporal resolution, boundary conditions, and prior error structure (spatial and temporal correlation scales), thus using ensembles of such systems is expected to result in more robust top-down estimates.
For this study, the global inversion results are taken from all six of the models reported in the GCP 2021:
CTE (CarbonTracker Europe), CAMS (Copernicus Atmosphere Monitoring Service), CMS-Flux, JENA, NIESNIWA, and UoE, with spatial resolutions ranging from 1°x1° for certain regions to 4°x5°. For details see Friedlingstein
et al. (2022). Note that one of the ensemble members (CMS-Flux) only covers the period 2010-2020, and therefore
the ensemble results are only shown from 2010 until the last year common between all models (2018).

2116 **EUROCOM**

2115

2117 Top-down estimates at regional scales (up to $0.25^{\circ}x0.25^{\circ}$ resolution) for the period 2009 – 2018 are taken 2118 from three models used within EUROCOM (Monteil et al., 2020): LUMIA, PYVAR, and CSR. The NAME model 2119 was excluded as visual inspection of monthly values identified it as a clear outlier. FLEXINVERT was excluded after 2120 visual inspection of annual values identified it as a clear outlier (Fig. A5). These inversions make use of more than 2121 30 atmospheric observing stations within Europe, including flask data and continuous observations. The CarboScope-2122 Regional (CSR) inversion system results were re-run for VERIFY using the extended period 2009-2020 using four 2123 different settings: three network configurations using 15, 40, or 46 sites, and one using all 46 sites but a factor two larger prior error correlation length scale (200 instead of 100 km). The CSR results reported to EUROCOM were not 2124 2125 used, being instead replaced by the mean of the four updated CSR runs.











CC VERIFY Project

Figure A5: Annual (top) and monthly (bottom) timeseries for inversions in EUROCOM (Monteil et al., 2020).
Inversions with solid lines were retained for the ensemble used in this work (shown in blue in the top figure for clarity).
Note that the CSR values from EUROCOM have been replaced by the mean of four CSR simulations submitted under
the VERIFY project (Appendix A). Negative fluxes represent a sink into the land surface.

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2134

2135 Input data

2136 CRUERA

The ERA5-Land (Muñoz-Sabater, 2019; 2021) dataset at 0.1-degree resolution over the global land surface 2137 2138 at hourly resolution was aggregated to three-hourly resolution and extracted for a 0.125 degree grid over Europe 2139 (35N:73N, 25W:45E) to match the grid used in previous efforts within the VERIFY project. The variables extracted 2140 are:air temperatures, wind components, surface pressure, downwelling longwave radiation, downwelling shortwave 2141 radiation, snowfall, and total precipitation. From these, additional variables were calculated: total windspeed, specific 2142 humidity, relative humidity, and rainfall. Of these, the air temperature, downwelling shortwave radiation, specific 2143 humidity, and total precipitation were re-aligned with the CRU observation dataset (Harris et al., 2020) from 1901-2144 2020 so that monthly means at 0.5 degree pixels correspond exactly. Variation from observations is therefore present 2145 only on sub-monthly temporal scales and sub-0.5 degree spatial scales. At the time of the model intercomparison, 2146 ERA5-Land was only available from 1981-2020. Consequently, the years 1901-1980 were taken from the UERRA 2147 HARMONIE-V1 dataset from ECMWF re-aligned with CRU observations under the VERIFY project and used in 2148 Petrescu et al. (2021b). For both datasets, results were aggregated to daily and monthly temporal resolution for use 2149 as needed in some models.

2150

2151 *HILDA*+

2152 The full Hilda+ dataset is described in detail elsewhere (Winkler et al., 2020; Winkler et al., 2021). Hilda+ 2153 is available at 1x1km spatial and annual temporal resolution across the whole globe from 1960-2019 for six land use 2154 classes (urban, cropland, pasture/rangeland, forest, unmanaged grass/shrubland, and sparse/no vegetation). The 2155 algorithm uses earth observation data and land use statistics to generate annual land use/cover maps and transitions. 2156 Probability maps for land use change categories are generated by using multiple earth-observation-based data 2157 estimates of the extent of a given land cover category on a given pixel. The VERIFY project requires additional work 2158 to satisfy the needs of the various modeling groups. For example, the maps were extended back to 1900 to meet the 2159 needs of the DGVM groups. As observational data is lacking for the years pre-1960, the temporal trend of the 2160 probability maps and the FAO land use database were used for extrapolation. In addition, forest areas were further 2161 subdivided into six forest types (Evergreen, needle leaf; Evergreen, broad leaf; Deciduous, needle leaf; Deciduous, 2162 broad leaf; Mixed; Unknown/Other) based on the ESA CCI land cover dataset (ESA 2017). Spatiotemporal forest type dynamics within the forest category were included for 1992-2015. Before 1992 and after 2015, the static forest type 2163 2164 distribution as found in the years 1992 and 2015 in the ESA CCI land cover was assumed, respectively.

2165

2166 NITROGEN DEPOSITION

Wet and dry deposition maps of ammonium and nitrate covering Europe from 1995-2018 were calculated at
0.5 degree spatial and monthly temporal resolution by the EMEP MSC-W model ("EMEP model" hereafter). The
EMEP model is a 3-D Eulerian chemistry transport model (CTM) developed at the EMEP Centre MSC-W under the
Framework of the UN Convention on Long-Range Transboundary Air Pollution (CLRTAP). The EMEP model has





2171 traditionally been used to assess acidification, eutrophication and air quality over Europe, to underpin air quality policy 2172 decisions (e.g., the Gothenburg Protocol), and has been under continuous development reflecting new scientific 2173 knowledge and increasing computer power. The model was described in detail by Simpson et al. (2012) and later 2174 updated as described in the annual EMEP status reports (Simpson et al., 2022, and references therein). For the 2175 VERIFY project, output from the EMEP model version rv4.33 was used (Simpson et al., 2019), and averaged to annual 2176 temporal resolution. In these simulations, the model was driven by meteorological data from the ECWMF IFS 2177 (European Centre for Medium-Range Weather Forecasts - Integrated Forecast System) version cy40r1. Land-use data 2178 were taken from the CORINE land-cover maps (de Smet and Hettelingh, 2001), the Stockholm Environment Institute 2179 at York (SEIY), the Global Land Cover (GLC2000) database, and the Community Land Model (Oleson et al., 2010; 2180 Lawrence et al., 2011). For more details see Simpson et al. (2017).

2181

2182 COASTAL OCEAN FLUXES

2183 Ocean CO₂ fluxes were prepared for use as prior estimates in the regional inversions by combining the Rödenbeck 2184 global ocean estimate (Rödenbeck et al., 2014) with coastal ocean fluxes for Europe prepared under the VERIFY 2185 project. The combined dataset was prepared by choosing the coastal flux map when available and otherwise the open 2186 ocean map. The coastal ocean fluxes were generated for an area extending from the western Mediterranean to the 2187 Barents Sea and cover shelf areas down to 500 m water depth or 100 km distance from shore. First, surface ocean 2188 fCO₂ observations are taken from the annually updated SOCAT database (Bakker et al., 2016) and gridded to a 2189 monthly 0.125° x0.125° grid. pCO₂ maps are created based on fitting a set of driver data (including sea surface 2190 temperature, mixed layer depth, chlorophyll concentration, and ice concentration) against the gridded fCO2 2191 observations. Both random forest and multi-linear regressions were used. The general procedure is described 2192 elsewhere (Becker et al., 2021), but for the version reported here, random forest regressions were used instead of 2193 multi-linear regression and the region was extended to the south. The dataset was divided into seven subregions (Barents Sea, Norwegian Coast, North Sea, Baltic Sea, Northern Atlantic Coast/Celtic Sea, Southern Atlantic 2194 2195 coast/Bay of Biscay, western Mediterranean) and each region was fitted separately (leaf size: 20, bag size: 500). The 2196 root mean square error (RMSE) of the random forest regressions was determined to be between 34 micro-atm (Baltic 2197 Sea) and 10 micro-atm (Barents Sea). Random forest regressions consist of many regression trees, each based on a 2198 random subset of data. Due to this internal structure, the overall RMSE can be seen as an out-of-box error estimate. 2199 The final fluxes are calculated from the pCO₂ maps with the atmospheric xCO₂ in the marine boundary layer and six-2200 hourly wind speed data using the gas transfer coefficient and the Schmidt number after Wanninkhoff (2014), the 2201 coefficient ag of 0.2814 calculated after Naegler (2009) and 6-hourly winds from the NCEP-DOE Reanalysis 2 product 2202 (Kanamitsu et al., 2002). 2203

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2205 Appendix B

2206 Overview figures



2207

Figure B1: EU27+UK total annual GHG emissions from UNFCCC NGHGI (2021) submissions split per
 sector.







Figure B2: EU27+UK total annual GHG emissions from the LULUCF sector split in categories and sub-categories, according to UNFCCC NGHGI (2021).

2214

2215 CO2 fossil breakdown by fuel type

Figure B3 shows the CO₂ fossil emission estimates from EU27+UK split by major source categories for each dataset for a single year. Sectors 1, 2, 3, and 5 are included for the UNFCCC NGHGI (2021) total, without indirect emissions. A breakdown of the nine other fossil BU data sources corresponding to UNFCCC NGHGI sectors or categories is not currently available.

As in Andrew (2020), we observe good agreement for the EU27+UK between all BU data sources and the UNFCCC NGHGI (2021) data. The figure presents updated estimates for the year 2017, the most recent year when all datasets reported estimates. Sectors 1, 2, 3, and 5 are included for the UNFCCC NGHGI (2021) total, without indirect emissions.

While most datasets agree well on total emissions, there are some differences. Both BP and the EIA include bunker fuels and exclude most industrial process emissions. CEDS appears to be underestimating emissions from solid fuels, for example lignite in Germany and oil shale in Estonia. IEA's emissions are lower because they exclude most industrial processes. GCP's total matches the NGHGIs exactly by design but remaps some of the fossil fuels used in non-energy processes from "Others" to the fuel types used. CDIAC, PRIMAP, and EDGAR v6.0 all report total emissions very similar to the UNFCCC NGHGI (2021). Larger differences are seen in the disaggregation of fuel types, generally because of differing definitions.



Comparison of fossil CO₂ emissions: EU-27+UK 2017





2232	Figure B3: EU27+UK total CO ₂ fossil emissions, as reported by nine bottom-up data sources: BP, EIA, CEDS, EDGAR v6.0, GCP, IEA,
2233	CDIAC, PRIMAPv2.3.1-CR and the UNFCCC NGHGI (2021) along with a top-down CIF-CHIMERE atmospheric inversion (black dot)
2234	(Fortems-Cheiney and Broquet, 2021). This figure presents the split per fuel type for year 2017. "Others" is other
2235	emissions in the UNFCCC's IPPU, and international bunker fuels (the white boxes) are not usually included in total
2236	emissions at sub-global level. Neither EDGAR ²⁸ (v6.0) nor PRIMAP publish a break-down by fuel type, so only the
2237	total is shown. For BP, the method description allows for emissions from natural gas to be calculated from BP's
2238	energy data, but the data for solid and liquid fuels are insufficiently disaggregated to allow replication of BP's
2239	emissions calculation method for those fuels.
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2243 Source specific methodologies: AD, EFs and uncertainties

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2245 Table B1: Source specific activity data (AD), emission factors (EF) and uncertainty methodology for all current 2246 VERIFY and non-VERIFY 2021 data products.

Data sources CO ₂ emission calculation	AD/Tier	EFs/Tier	Uncertainty assessment method	Emission data availability
UNFCCC NGHGI (2021)	Country-specific information consistent with the IPCC Guidelines	IPCC guidelines / Country specific information for higher Tiers	IPCC guidelines (https://www.ipcc- nggip.iges.or.jp/public/2006g I/) for calculating the uncertainty of emissions based on the uncertainty of AD and EF, two different approaches: 1. Error propagation, 2. Monte Carlo Simulation UBA Vienna provided yearly harmonized and gap-filled uncertainties	NGHGI official data (CRFs) are found at https://unfccc.int/ghg- inventories-annex-i-parties/2021 (last access: June 2022).

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Fossil CO₂

²⁸EDGAR v6.0 provides significant sectoral disaggregation of emissions, but not by fuel type due to license restrictions with the underlying energy data from the IEA.





BP CDIAC EIA IEA GCP CEDS PRIMAP-Hist	For further details, see Andrew (2020)									
EDGAR v6.0	International Energy Agency (IEA) for fuel combustion Food and Agricultural Organisation (FAO) for agriculture US Geological Survey (USGS) for industrial processes (e.g., cement, lime, ammonia and ferroalloys production) GGFR/NOAA for gas flaring World Steel Association for iron and steel production International Fertilisers Association (IFA) for urea consumption and production Complete description of the data sources can be found in Janssens- Maenhout et al. (2019) and in Crippa et al. (2019)	IPCC (2006): Tier 1 or Tier 2 depending on the sector	Tier 1 with error propagation by fuel type for CO ₂ and accounting for covariances.	https://edgar.jrc.ec.europa.eu/datas et_ghg60						
CIF- CHIMERE	Tier 3 top-down 0.1° x 0.1° resolution maps of annual averages of fossil CO ₂ anthropogenic emissions from EDGAR v4.3.2 Assimilation of satellite atmospheric concentration data: total column CO from IASI, and tropospheric column NO ₂ from OMI	Tier 3 top-down regional inversions of CO and NOx emissions using EMEP/CEIP as prior knowledge of the emissions and CO ₂ /CO and CO ₂ /NOx emission ratios associated with the combustion of fossil fuel from EDGARv4.3.2.	Bayesian analysis in the CO and NOx inversions along with propagation of uncertainties in fCO ₂ /CO and fCO ₂ /NOx emission ratios	Detailed gridded data can be obtained by contacting the data providers: Gregoire Broquet gregoire.broquet@lsce.ipsl.fr						

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CO2 land: bottom-up





BLUEVGCP	From LUH2: data on wood harvest, land cover types (primary, secondary, pasture, crop), and gross land use transitions (e.g. from secondary to pasture and back); Based on Pongratz et al. (2008) and Ramankutty and Foley (1999): Plant functional types (PFTs) of natural vegetation types	Tier 3 (IPCC, 2006); PFT and land-cover type specific response curves describing the decay and regrowth of vegetation and soil carbon	N/A	Detailed gridded data can be obtained by contacting the data provider: Julia Pongratz: julia.pongratz@lmu.de
BLUEVVERIF Y	Same as above with land cover from HILDA+ (Ganzenmüller et al., 2022)			
H&N	Simple assumptions about C-stock densities (per biome or per biome/country) based on literature	Transient change in C-stocks following a given transition (time dependent EF after an land use transition)	N/A	Detailed gridded data can be obtained by contacting the data provider: Richard A. Houghton rhoughton@woodwellclimate.org
ECOSSE	Tier 3 approach. The model is a point model, which provides spatial results by using spatial distributed input data (lateral fluxes are not considered). The model is a Tier 3 approach that is applied on grid map data, polygon organized input data or study sites.	IPCC (2006): Tier 3 The simulation results will be allocated due to the available information (size of spatial unit, representation of considered land use, etc.).	N/A	Detailed gridded data can be obtained by contacting the data providers: Kuhnert, Matthias <u>matthias.kuhnert@abdn.ac.uk</u> Pete Smith: pete.smith@abdn.ac.uk
EPIC-IIASA Croplands	Tier 3 approach. Cropland: static 1×1 km cropland mask from CORINE-PELCOM. Initial SOC stock from the Map of organic carbon content in the topsoil (Lugato et al., 2014). "Static" crop management and input intensity by NUTS2 calibrated for 1995- 2010 (Balkovič et al., 2013). Crop harvested areas by NUTS2 from EUROSTAT. Parameterization of soil carbon routine was	IPCC (2006): Tier 3 Land management and input factors for the cropland remaining cropland category as simulated by the EPIC-IIASA modeling platform, assuming the business-as-usual crop management calibrated for the 1995-2010	Sensitivity and uncertainty analysis of EPIC-IIASA regional soil carbon modeling (Balkovič et al, 2020).	Detailed gridded data can be obtained by contacting the data provider: Balcovič Juraj balkovic@iiasa.ac.at



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	updated based on Balkovič et al. (2020).	period. A 50-ha field is considered in each grid cell.		
EPIC-IIASA grasslands	Tier 3 approach. Grassland: static 1x1 km mask from CORINE & PELCOM 2000, including pastures, herbaceous vegetation, heterogeneous agricultural areas, and permanent cropland. Initial SOC stock from the map of organic carbon content in the topsoil (Lugato et al., 2014) with a spin-up. Static grassland management and input intensity as adopted from (Chang et al., 2016) and ISIMIP (Jägermeyr et al., 2021).	IPCC (2006): Tier 3 Land management and input factors for the grassland remaining grassland category as simulated by the EPIC-IIASA modeling platform, calibrated for the 1995–2020 period.	N/A	Detailed gridded data can be obtained by contacting the data provider: Juraj Balkovič: balkovic@iiasa.ac.at
ORCHIDEE	For the land cover/land use input maps: data on wood harvest from the FAO	Tier 3 model, process based. Any emission factors enter in the form of generic parameters for a given ecosystem type fit against observational data (both site- level and remotely sensed).	None, though some information on uncertainty due to model structure is given by looking at the spread from the TRENDY suite of models, of which ORCHIDEE is a member.	Detailed gridded data can be obtained by contacting the data providers: Matthew McGrath <u>matthew.mcgrath@lsce.ipsl.fr</u> Philippe Peylin: <u>peylin@lsce.ipsl.fr</u>
CABLE-POP	For the land cover/land use input maps: data on wood harvest and agricultural land from the FAO	Tier 3 model, process based. Any emission factors enter in the form of generic parameters for a given ecosystem type fit against	None, though some information on uncertainty due to model structure is given by looking at the spread from the TRENDY suite of models, of which CABLE-POP is a member.	Model output (gridded data) can be obtained by contacting the data provider: Jürgen Knauer: J.Knauer@westernsydney.edu.au





			observational data (both site- level and remotely sensed).		
2250	TRENDY v10	For the land cover/land use input maps: data on wood harvest and agricultural land from the FAO	Tier 3 models, process based. Any emission factors enter in the form of generic parameters for a given ecosystem type fit against observational data (both site- level and remotely sensed).	The spread of the 15 TRENDY models used gives an idea of the uncertainty due to model structure in dynamic global vegetation models, as the forcing data was harmonized for all models.	Detailed gridded data can be obtained by contacting the data provider: Sitch, Stephen S.A.Sitch@exeter.ac.uk
	Statistical prediction model for CO ₂ in inland waters	Hydrosheds 15s (Lehner et al., 2008) and Hydro1K (USGS, 2000) for river network, HYDROLAKES for lakes and reservoirs network and surface area (Messager et al., 2016); river pCO2 data from GloRiCh (Hartmann et al., 2014), lake pCO ₂ database from Sobek et al. (2005); river channel slope and width calculated from GLOBE-DEM (GLOBE-Task-Team et al., 1999) and runoff data from Fekete et al. (2002). Geodata for predictors of pCO ₂ and gas transfer coefficient include air temperature, precipitation and wind speed (Hijmans et al., 2005), population density (CIESIN and CIAT), catchment slope gradient (Hydrosheds 15s), and terrestrial NPP (Zhao et al., 2005)	N/A	Monte Carlo runs (uncertainty on pCO ₂ and gas transfer velocity)	Detailed gridded data can be obtained by contacting the data providers: Ronny Lauerwald <u>Ronny Lauerwald@ulb.ac.be</u> Pierre Regnier Pierre.Regnier@ulb.ac.be
	СВМ	National forest inventory data, Tier 2	EFs directly calculated by model, based on specific parameters (i.e., turnover and decay rates)	N/A used from IPCC	Detailed gridded data can be obtained by contacting the data providers: Giacomo Grassi <u>Giacomo.GRASSI@ec.europa.eu</u> Matteo Vizzarri <u>Matteo.VIZZARRI@ec.europa.eu</u>





		defined by the user		Roberto Pilli roberto.pilli713@gmail.com
EFISCEN- Space	National forest inventory data, Tier 3	emission factor is calculated from net balance of growth minus harvest	Sensitivity analysis on EFISCEN V3 in the user manual (Schelhaas et al., 2007). Total sensitivity is caused by esp. young forest growth, width of volume classes, age of felling and few more. Scenario uncertainty comes on top of this when projecting in future.	Detailed gridded data can be obtained by contacting the data providers: Gert-Jan Nabuurs <u>gert-jan nabuurs@wur.nl</u> Mart-Jan Schelhaas <u>martjan.schelhaas@wur.nl</u>
FAOSTAT	FAOSTAT Land Use Domain; Harmonized world soil; ESA CCI; MODIS 6 Burned area products	IPCC guidelines	IPCC (2006, Vol.4, p.10.33) - confidential Uncertainties in estimates of GHG emissions are due to uncertainties in emission factors and activity data. They may be related to, inter alia, natural variability, partitioning fractions, lack of spatial or temporal coverage, or spatial aggregation.	Agriculture total and subdomain specific GHG emissions are found for download at http://www.fao.org/faostat/en/#dat a/GT (last access: April 2022).
		CO ₂ land	: Top-down	
CSR GCP ensemble (CTE, CAMS, CarboScope) EUROCOM (PYVAR- CHIMERE, LUMIA, FLEXINVERT, CSR, CTE- Europe) LUMIA CIF- CHIMERE	Tier 3 top-down approach, prior information from fossil emissions, ocean fluxes, and biosphere- atmosphere exchange Spatial resolutions ranging from 1°x1° for certain regions to 4°x5°. EUROCOM uses more than 30 atmospheric stations. CSR uses four different settings (as described in Appendix A2)	Tier 3 top-down Inversion systems based on atmospheric transport models	$\label{eq:construction} \begin{aligned} \mathbf{CSR} & - \text{Gaussian probability} \\ \text{distribution function, where} \\ \text{the error covariance matrix} \\ \text{includes errors in prior} \\ \text{fluxes, observations and} \\ \text{transport model} \\ \text{representations.} \\ \mathbf{GCP}: \text{the different} \\ \text{methodologies, the land-use} \\ \text{and land-cover data set, and} \\ \text{the different processes} \\ \text{represented trigger the} \\ \text{uncertainties between} \\ \text{models. a semi-quantitative} \\ \text{measure of uncertainty for} \\ \text{annual and decadal} \\ \text{emissions as best value} \\ \text{judgment} = \text{at least a 68\%} \\ \text{chance } (\pm 1\sigma) \\ \textbf{EUROCOM: account for} \\ \text{source of uncertainties via} \\ \text{prior and model and} \\ \text{observation error covariance} \\ \text{matrices; assessment of the} \\ \text{resulting uncertainties in} \\ \text{fluxes based on spread} \\ \textbf{LUMIA:} \\ \\ \text{The prior uncertainties are} \\ \text{constructed using standard} \\ \end{array}$	Detailed gridded data can be obtained by contacting the data providers: CSR: Christoph Gerbig cgerbig@bgc- jena.mpg.de Saqr Munassar smunas@bgc- jena.mpg.de GCP: Pierre Friedlingstein P.Friedlingstein@exeter.ac.uk EUROCOM: Marko Scholze marko.scholze@nateko.lu.se Gregoire Broquet gregoire.broquet@lsce.ipsl.fr LUMIA: Guillaume Monteil guillaume.monteil@nateko.lu.se CIF-CHIMERE: Gregoire Broquet gbroquet@lsce.ipsl.fr





	deviations proportional to the sum of the absolute value of the hourly NEE aggregated in each weekly optimization interval (so, in essence, uncertainties are large when the daily cycle of NEE is large), spatial correlation lengths of 500 km (Gaussian) and temporal correlation lengths of 1 month (Exponential).	
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2253	Table B2:	Comparison	of the processes	included in the invent	ories, bottom-up mod	els and inversions.
		1	v 1			

Descri- ption	NGHGI	Global database	Process-based models			DGVMs			Bookkeeping Models			Inversions [#]	
	U N F C C C	F A O S T A T ^a	E C S S E	E P I C - I I A S A	C B M	E F I S C E N - Space	C A B L E - P O P	T R E N D Y V 1 0	O R C H I D E E	B L U E v G C P	B L U E V E R I F Y	H & N	
Forest total	Е	Е	Ν	Ν	Е	Е	Е	Acc. table A1 in GCB	Е	E^{h}	E^{h}	E^{h}	
Split FL-FL / FL-X / X-FL	E	Е	Ν	Ν	Е	E/N/N	Е	2021 (Friedling stein et al., 2022)	Е	E ^h /E/	E ^h /E/	E ^h /E/ E	
Croplan d total	Е	N	Е	Е	N	N	Ι		Е	E^{h}	E^{h}	E^{h}	
Split CL-CL / CL-X / X-CL	E	Ν	E	E/N/ N	N	Ν	I		Е	N/E/E	N/E/E	N/E/E	
Grasslan d total	Е	Ν	Е	N	Ν	Ν	Е		Е	Е	Е	Е	
Split GL-GL / GL-X / X-GL	Е	Ν	Е	N	N	Ν	Е		Е	N/E/E	N/E/E	N/E/E	
Peatland accounti ng	Е	E	N	N	N	N	N		N	N	Ν	N	





CO ₂ fertilizat ion	Ι	Ι	Ν	Е	Ν	N	Е	Acc. table A1 in GCB	Е	N ⁱ	N ⁱ	N ⁱ	
Climate induced impacts	Ι	Ι	Ν	E ^f	Ip	Ic	Е	(Friedling stein et al., 2022)	Е	Ni	N ⁱ	N ⁱ	
Natural disturba nces (fires, insect, wind)	Ι	Ι	N	N	Е	N	E		N	N ⁱ	N ⁱ	N ⁱ	
Soil Organic C dynamic	I		Е	Е	Е	Е	E		Е	N	N	N	
Lateral C transport (river)	N	Ν	Ν	Ν	Ν	Ν	N		Ν	N	N	N	
Flux from Harveste d Wood Products	Е	Ν	N	Ν	Ι	N ^d	E	Acc. table A1 in GCB 2021 (Friedling	Е	Е	Е	Е	
Flux from Crop/Gr ass harvest	?	Ν	Е	Ee	Ν	Ν	Е	al., 2022)	Е	Ii	I	Ii	
Biomass burning	Е	Е	Е	N ^g	Е	Ν	Ν		N	Ej	Ej	Ej	
N fertilizat ion (with N dep)	I	N	Е	Ν	Ν	Ν	Е		N	N	N	N	
Flux from drained organic soils	I	Е	Е	N	Ι	N	N		Ι	Ej	Ej	Ej	
2254 2255 2256 2257	^a UNFCCC a ^b The climate	Not included : and FAOSTA e effects can b	N, Expl T are ens e estima	icitly mo emble o ted indir	deled : I f country ectly by	E, Implicitly m v estimates cal CBM, using e	odeled: I, I culated with xternal add	Partly modeled h specific meth itional input p	l : P nodology rovided l	for each	country, f 10dels	following s	some guidelines

°EFISCEN Space: Increment is sensitive to weather, but average weather

dEFISCEN has only production in m3 but doesn't have a direct HWP module

^eCrop yield and residue harvest from cropland (20 % of residues harvested in case of cereals, no residue harvest for other crops)

EPIC-IIASA partly accounts for soil drought, i.e., plant growth limitation due to a lack of water in the soils. Heat stress and floods are not accounted for, though

gIn principle, burning of crop residues on cropland can be explicitly simulated by EPIC-IIASA. However, not done for VERIFY as it is not a relevant scenario for the business as usual cropland management in Europe

2258 2259 2260 2261 2262 2263 2264 2265 2266 2265 2266 2267 2268 2269 hforest/cropland/grassland exist and have carbon stocks, but have carbon fluxes only through change to management. FL-FL includes all land-use induced effects (harvest slash and product decay, regrowth after agric abandonment and harvesting)

'implicit by using observation-based carbon densities that reflect harvest/climate/natural disturbances

peat burning and peat drainage are not bookkeeping model output, but are added from various data sources during post processing

*According Table 2 in Monteil et al. (2020) and Table A3 in Friedlingstein et al. (2019)





2270 **These categories are inputs to the inversions, not a result; the inversions adjust the total land-atmosphere C flux, regardless of what went into the prior, and the posterior flux cannot really be disaggregated into contributions from separate processes. In a sense, as long as a process is sufficiently significant to influence the CO₂ observations, it will have an impact on the inversion results

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2276 Author contributions

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2278 MJM processed original data, made Fig. 1,3-10, A2, A3, A5, and edited the final manuscript; AMRP designed the initial research, led the discussions, wrote the initial draft of the paper and helped edit all the following versions; RMA 2279 2280 made Fig. 2, A1, B3; BM provided the new UNFCCC gap-filled uncertainties and provided extensive support on 2281 questions related to NGHGIs; PP, VB, and MJM processed the original data submitted to the VERIFY portal; PP, PB, 2282 and MJM designed and are managing the web portal; GP provided Fig. B1 and B2; GP, RMA, FD, BM, and GG made 2283 detailed reviews; CQ made Fig. 11; SM made Fig. A4; PC, GB, PIP, MJ, RL, MK, JK, FC, OT, JP, RG, FNT, JB and 2284 GG gave detailed comments and advice on previous versions of the manuscript; all remaining co-authors provided 2285 data and commented on specific parts of the text related to their data sets.

2286

2287 Competing interests

2288 The authors declare that they have no conflict of interest.

2289

2290 Acknowledgements

2291 We thank Aurélie Paquirissamy, Géraud Moulas and all ARTTIC team, for the great managerial support offered during 2292 the VERIFY project. FAOSTAT statistics are produced and disseminated with the support of its member countries 2293 to the FAO regular budget. The views expressed in this publication are those of the author(s) and do not necessarily 2294 reflect the views or policies of FAO. We acknowledge the work of other members of the EDGAR group (Edwin 2295 Schaaf, Jos Olivier). We acknowledge Stephen Sitch and the authors of the DGVMs TRENDY v10 ensemble models 2296 for providing us with the data. We thank all the national forest inventories that have made their data available: Ireland 2297 (John Redmond), Norway (Rasmus Astrup), Sweden (Jonas Fridman), Poland (Andrzej Talarczyk), Germany 2298 (BMEL), The Netherlands (WUR & Stichting Probos), Belgium (Flanders: Leen Govaere), Luxembourg (Thierry 2299 Palgen), France (IGN), Spain (MAPA), Switzerland (Esther Thürig), Italy (CREA), Czech Republic (Emil Cienciala), 2300 Slovak Republic (Vladimír Šebeň). We thank all the NFI field crews for their hard work. Timo Vesala thanks ICOS-2301 Finland, University of Helsinki. Ingrid T. Luijkx and Wouter Peters thank the HPC cluster Aether at the University of 2302 Bremen, financed by DFG within the scope of the Excellence Initiative. MJM and VB were granted access to the 2303 HPC resources of GENCI-TGCC under the allocation A0130106328. 2304

2305

2306 Financial support

This research has been supported by the European Commission, Horizon 2020 Framework Programme (VERIFY,
grant no. 776810, for AB, AFC, AMRP, AP, CG, GB, GJM, GJN, GM, GP, HACDG, JB, LP, MJ, MJM, MK, MV,

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- 2309 PP, PR, PS, RG, RMA, SD). MJM and GM also acknowledge funding from the European Union's Horizon 2020
- 2310 research and innovation programme under Grant Agreement No. 958927 (CoCO2). Philippe Ciais acknowledges
- 2311 the support of European Research Council Synergy project SyG-2013-610028 IMBALANCE-P and from the ANR
- 2312 CLand Convergence Institute. Ronny Lauerwald thanks the CLand Convergence Institute. Pierre Regnier
- acknowledges the ESM 2025. Gert-Jan Nabuurs thanks the Dutch National Forest Inventory funded by the Ministry
- 2314 Agriculture Nature Management and Food Quality. Guillaume Monteil's model computations were enabled by
- 2315 resources provided by the Swedish National Infrastructure for Computing (SNIC) at NSC partially funded by the
- 2316 Swedish Research Council through grant agreement no. 2018-05973.





2317	References
2318	Andrew, R. M.: A comparison of estimates of global carbon dioxide emissions from fossil carbon sources, Earth Syst.
2319	Sci. Data, 12, 1437–1465, https://doi.org/10.5194/essd-12-1437-2020, 2020.
2320	
2321	Arneth, A., Sitch, S., Pongratz, J., Stocker, B. D., Ciais, P., Poulter, B., Bayer, A. D., Bondeau, A., Calle, L., Chini,
2322	L. P., Gasser, T., Fader, M., Friedlingstein, P., Kato, E., Li, W., Lindeskog, M., Nabel, J. E. M. S., Pugh, T. A. M.,
2323	Robertson, E., Viovy, N., Yue, C., and Zaehle, S.: Historical carbon dioxide emissions caused by land-use changes
2324	are possibly larger than assumed. Nature Geosci, 10, 79-84, https://doi.org/10.1038/ngeo2882, 2017.
2325	
2326	Balkovič, J., Madaras, M., Skalský, R., Folberth, C., Smatanová, M., Schmid, E., van der Velde, M., Kraxner, F.,
2327	Obersteiner, M.: Verifiable soil organic carbon modeling to facilitate regional reporting of cropland carbon change:
2328	A test case in the Czech Republic, J. Environ. Manage., 274, 111206, https://doi.org/10.1016/j.jenvman.2020.111206,
2329	2020.
2330	
2331	Balkovič, J., Skalský, R., Folberth, C., Khabarov, N., Schmid, E., Madaras, M., Obersteiner, M., van der Velde, M.:
2332	Impacts and Uncertainties of +2°C of Climate Change and Soil Degradation on European Crop Calorie Supply, Earths
2333	Future, 6, 373–395, https://doi.org/10.1002/2017EF000629, 2018.
2334	
2335	Balkovič, J., van der Velde, M., Schmid, E., Skalský, R., Khabarov, N., Obersteiner, M., Stürmer, B., Xiong, W.: Pan-
2336	European crop modeling with EPIC: Implementation, up-scaling and regional crop yield validation, Agric. Syst., 120,
2337	61-75, https://doi.org/10.1016/j.agsy.2013.05.008, 2013.
2338	
2339	Bastos, A., Ciais, P., Friedlingstein, P., Sitch, S., Pongratz, J., Fan, L., Wigneron, J. P., Weber, U., Reichstein, M., Fu,
2340	Z., Anthoni, P., Arneth, A., Haverd, V., Jain, A. K., Joetzjer, E., Knauer, J., Lienert, S., Loughran, T., McGuire, P. C.,
2341	Tian, H., Viovy, N., and Zaehle, S.: Direct and seasonal legacy effects of the 2018 heat wave and drought on European
2342	ecosystem productivity, Science Advances, 6, eaba2724, DOI: 10.1126/sciadv.aba27, 2020a.
2343	



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Pongratz, J., Luijkx, I. T., Patra, P. K., Peylin, P., Canadell, J. G., Lauerwald, R., Li, W., Smith, N. E., Peters, W.,							
Goll, D. S., Jain, A. K., Kato, E., Lienert, S., Lombardozzi, D. L., Haverd, V., Nabel, J. E. M. S., Poulter, B., Tian,							
H., Walker, A. P., and Zaehle, S.: Sources of uncertainty in regional and global terrestrial CO 2 exchange estimates,							
Global Biogeochemical Cycles, 34, e2019GB006393, https://doi.org/10.1029/2019GB006393, 2020b.							
Becker, M., Olsen, A., Landschützer, P., Omar, A., Rehder, G., Rödenbeck, C., and Skjelvan, I.: The northern							
European shelf as an increasing net sink for CO2, Biogeosciences, 18, 1127-1147, https://doi.org/10.5194/bg-18-							
1127-2021, 2021.							
Berchet, A., Sollum, E., Thompson, R. L., Pison, I., Thanwerdas, J., Broquet, G., Chevallier, F., Aalto, T., Berchet,							
A., Bergamaschi, P., Brunner, D., Engelen, R., Fortems-Cheiney, A., Gerbig, C., Groot Zwaaftink, C. D., Haussaire,							
JM., Henne, S., Houweling, S., Karstens, U., Kutsch, W. L., Luijkx, I. T., Monteil, G., Palmer, P. I., van Peet, J. C.							
A., Peters, W., Peylin, P., Potier, E., Rödenbeck, C., Saunois, M., Scholze, M., Tsuruta, A., and Zhao, Y.: The							
Community Inversion Framework v1.0: a unified system for atmospheric inversion studies, Geosci. Model Dev., 14,							
5331-5354, https://doi.org/10.5194/gmd-14-5331-2021, 2021.							
BP: 60 Years BP Statistical Review of World Energy: 1951-2011, available at:							
$https://www.bp.com/en/global/corporate/\ energy-economics/statistical-review-of-world-energy/downloads.html\ (last of the second seco$							
access: 8 February 2019), 2011.							
BP: BP Statistical Review of World Energy June 2018, available at: https://www.bp.com/en/global/corporate/energy-							
economics/statistical-review-of-world-energy/downloads.html, (last access: 14 June 2018).							

Bastos, A., O'Sullivan, M., Ciais, P., Makowski, D., Sitch, S., Friedlingstein, P., Chevallier, F., Rödenbeck, C.,

2367

2368 BP: Methodology for calculating CO2 emissions from energy use, available at: 2369 https://www.bp.com/en/global/corporate/ energy-economics/statistical-review-of-world-energy/co2-emissions.html 2370 (last access: 8 February 2019), 2017.

<sup>Bradbury, N. J., Whitmore, A. P., Hart, P. B. S., and Jenkinson, D. S.: Modelling the fate of nitrogen in crop and soil
in the years following application of 15N-labelled fertilizer to winter wheat, J. Agr. Sci., 121, 363-379,
doi:10.1017/S0021859600085567, 1993.</sup>





2375	
2376	Brophy, K., Graven, H., Manning, A. J., White, E., Arnold, T., Fischer, M. L., Jeong, S., Cui, X., and Rigby, M.:
2377	Characterizing uncertainties in atmospheric inversions of fossil fuel CO2 emissions in California, Atmos. Chem.
2378	Phys., 19, 2991-3006, https://doi.org/10.5194/acp-19-2991-2019, 2019.
2379	
2380	Broquet, G., Chevallier, F., Rayner, P., Aulagnier, C., Pison, I., Ramonet, M., Schmidt, M., Vermeulen, A. T., and
2381	Ciais, P.: A European summertime CO2 biogenic flux inversion at mesoscale from continuous in situ mixing ratio
2382	measurements,
2383	J. Geophys. Res., 116, D23303, doi:10.1029/2011JD016202, 2011.
2384	
2385	Broquet, G., Chevallier, F., Bréon, FM., Kadygrov, N., Alemanno, M., Apadula, F., Hammer, S., Haszpra, L.,
2386	Meinhardt, F., Morguí, J. A., Necki, J., Piacentino, S., Ramonet, M., Schmidt, M., Thompson, R. L., Vermeulen, A.
2387	T., Yver, C., and Ciais, P.: Regional inversion of CO2 ecosystem fluxes from atmospheric measurements: reliability
2388	of the uncertainty estimates, Atmos. Chem. Phys., 13, 9039–9056, https://doi.org/10.5194/acp-13-9039-2013, 2013.
2389	
2390	CDIAC, https://energy.appstate.edu/CDIAC (last access: 10 November 2022).
2391	
2392	Ceccherini, G, Duveiller, G, Grassi, G, Lemoine, G, Avitabile, V, Pilli, R, and Cescatti, A.: Abrupt increase in
2393	harvested forest area over Europe after 2015, Nature, 583, 72–77, https://doi.org/10.1038/s41586-020-2438-y, 2020.
2394	
2395	CEDS v_2019_12_23, https://www.pnnl.gov/projects/ceds (last access: 10 November 2022).
2396	
2397	Chang, J., Ciais, P., Herrero, M., Havlik, P., Campioli, M., Zhang, X., Bai, Y., Viovy, N., Joiner, J., Wang, X., Peng,
2398	S., Yue, C., Piao, S., Wang, T., Hauglustaine, D. A., Soussana, JF., Peregon, A., Kosykh, N., and Mironycheva-
2399	Tokareva, N.: Combining livestock production information in a process-based vegetation model to reconstruct the
2400	history of grassland management, Biogeosciences, 13, 3757-3776, https://doi.org/10.5194/bg-13-3757-2016, 2016.
2401	
2402	Chevallier, F., Fisher, M., Peylin, P., Serrar, S., Bousquet, P., Bréon, FM., Chédin, A., and Ciais, P.: Inferring CO2
2403	sources and sinks from satellite observations: Method and application to TOVS data, J. Geophys. Res., 110, D24309,
2404	doi:10.1029/2005JD006390, 2005.

2405





- 2406 Chevallier, F., F.-M. Bréon, F.-M., and Rayner, P. J.: Contribution of the Orbiting Carbon Observatory to the2407 estimation of CO2 sources and sinks: Theoretical study in a variational data assimilation framework, J. Geophys. Res.,
- 2408 112, D09307, doi:10.1029/2006JD007375, 2007.
- 2409
- 2410 Chevallier, F., Ciais, P., Conway, T. J., Aalto, T., Anderson, B. E., Bousquet, P., Brunke, E. G., Ciattaglia, L., Esaki,
- 2411 Y., Fröhlich, M., Gomez, A. J., Gomez-Pelaez, A. J., Haszpra, L., Krummel, P., Langenfelds, R., Leuenberger, M.,
- 2412 Machida, T., Maignan, F., Matsueda, H., Morguí, J. A., Mukai, H., Nakazawa, T., Peylin, P., Ramonet, M., Rivier, L.,
- 2413 Sawa, Y., Schmidt, M., Steele, P., Vay, S. A., Vermeulen, A. T., Wofsy, S., and Worthy, D. : CO2 surface fluxes at
- 2414 grid point scale estimated from a global 21-year reanalysis of atmospheric measurements. J. Geophys. Res., 115,
- 2415 D21307, doi:10.1029/2010JD013887, 2010.
- 2416
- 2417 Ciais, P., Reichstein, M., Viovy, N., Granier, A., Ogée, J., Allard, V., Aubinet, M., Buchmann, N., Bernhofer Chr.,
- 2418 Carrara, A., Chevallier, F., De Noblet, N., Friend, A. D., Friedlingstein, P., Grünwald, T., Heinesch, B., Keronen, P.,
- 2419 Knohl, A., Krinner, G., Loustau, D., Manca, G., Matteucci, G., Miglietta, F., Ourcival, J. M., Papale, D., Pilegaard,
- 2420 K., Rambal, S., Seufert, G., Soussana, J. F., Sanz, M. J., Schulze, E. D., Vesala, T. and Valentini, R.: Europe-wide
- 2421 reduction in primary productivity caused by the heat and drought in 2003. Nature, 437, 529–533,
- 2422 https://doi.org/10.1038/nature03972, 2005.
- 2423
- 2424 Ciais, P., Crisp, D., Denier van der Gon, H., Engelen, R., Janssens-Maenhout, G., Heimann, M., Rayner, P., and
- 2425 Scholze, M.: Towards a European Operational Observing System to Monitor Fossil CO2 emissions Final Report
- from the expert group, https://www.copernicus.eu/sites/default/files/2019-09/CO2 Blue report 2015.pdf, 2015.
- 2427
- 2428 Ciais, P., Yao, Y., Gasser, T., Baccini, A., Wang, Y., Lauerwald, R., Peng, S., Bastos, A., Li, W., Raymond, P.A. and
- 2429 Canadell, J.G., Peters, G. P., Andres, R. J., Chang, J., Yue, C., Dolman, A. J., Haverd, V., Hartmann, J., Laruelle, G.,
- 2430 Konings, A. G., King, A. W., Liu, Y., Luyssaert, S., Maignan, F., Patra, P. K., Peregon, A., Regnier, P., Pongratz, J.,
- 2431 Poulter, B., Shvidenko, A., Valentini, R., Wang, R., Brouquet, G., Yin, Y., Zscheischler, J., Guenet, B., Goll, D. S.,
- 2432 Ballantyne, A.-P., Yang, H., Qiu, C., and Zhu, D.: Empirical estimates of regional carbon budgets imply reduced
- 2433 global soil heterotrophic respiration, National Science Review, 8, nwaa145, https://doi.org/10.1093/nsr/nwaa145,
- **2434** 2021.
- 2435
- 2436 CoCO2: https://coco2-project.eu/, last access: 21 November 2022.





- 2438 Coleman, K., Jenkinson, D. S.: RothC-26.3 A model the turnover of carbon in soil. In: Powlson DS, Smith P, Smith
- 2439 JU (ed) Evaluation of soil organic matter models using existing long-term datasets, NATO ASI Series I, vol. 38.
- 2440 Springer, Berlin, pp 237–246, 1996.
- 2441
- 2442 Crippa, M., Oreggioni, G., Guizzardi, D., Muntean, M., Schaaf, E., Lo Vullo, E., Solazzo, E., Monforti-Ferrario, F.,
- 2443 Olivier, J.G.J., and Vignati, E.: Fossil CO2 and GHG emissions of all world countries 2019 Report, EUR 29849 EN,
- 2444 Publications Office of the European Union, Luxembourg, 2019, ISBN 978-92-76-11100-9, doi:10.2760/687800,
- **2445** JRC117610, 2019.
- 2446
- Deng, Z., Ciais, P., Tzompa-Sosa, Z. A., Saunois, M., Qiu, C., Tan, C., Sun, T., Ke, P., Cui, Y., Tanaka, K., Lin, X.,
 Thompson, R. L., Tian, H., Yao, Y., Huang, Y., Lauerwald, R., Jain, A. K., Xu, X., Bastos, A., Sitch, S., Palmer, P.
 I., Lauvaux, T., d'Aspremont, A., Giron, C., Benoit, A., Poulter, B., Chang, J., Petrescu, A. M. R., Davis, S. J., Liu,
 Z., Grassi, G., Albergel, C., Tubiello, F. N., Perugini, L., Peters, W., and Chevallier, F.: Comparing national
 greenhouse gas budgets reported in UNFCCC inventories against atmospheric inversions, Earth Syst. Sci. Data, 14,
- 2452 1639–1675, https://doi.org/10.5194/essd-14-1639-2022, 2022.
- 2453
- Ducoudré, N. I., Laval, K., and Perrier, A.: SECHIBA, a new set of parameterizations of the hydrologic exchanges at
 the land-atmosphere interface within the LMD atmospheric general circulation model, Journal of Climate, 6, 248–
 273, https://www.jstor.org/stable/26197219, 1993.
- 2457
- ESA: Land Cover CCI Product User Guide Version 2. ESA. http://maps.elie.ucl.ac.be/CCI/viewer/index.php, (last
 access: 10 November 2022), 2017.
- 2460
- EU: REGULATION (EU) No 525/2013 OF THE EUROPEAN PARLIAMENT AND OF THE COUNCIL of 21 May
 2013 on a mechanism for monitoring and reporting greenhouse gas emissions and for reporting other information at
 national and Union level relevant to climate change and repealing Decision No 280/2004/EC, https://eur lex.europa.eu/legal-content/EN/TXT/PDF/?uri=CELEX:32013R0525&from=EN, 2013.
- 2465





2466	EU: Regulation (EU) 2018/841 of the European Parliament and of the Council of 30 May 2018 on the inclusion of
2467	greenhouse gas emissions and removals from land use, land use change and forestry in the 2030 climate and energy
2468	framework, and amending Regulation (EU) No 525/2013 and Decision No 529/2013/EU, https://eur-
2469	lex.europa.eu/legal-content/EN/TXT/?uri=uriserv:OJ.L2018.156.01.0001.01.ENG, 2018a.
2470	
2471	EU: Regulation (EU) 2018/842 of the European Parliament and of the Council of 30 May 2018 on binding annual
2472	greenhouse gas emission reductions by Member States from 2021 to 2030 contributing to climate action to meet
2473	commitments under the Paris Agreement and amending Regulation (EU) No 525/2013, https://eur-
2474	lex.europa.eu/legal-content/EN/TXT/?uri=CELEX:32018R0842, 2018b.
2475	
2476	EU: Communication COM/2020/562: Stepping up Europe's 2030 climate ambition Investing in a climate-neutral
2477	$future \ for \ the \ benefit \ of \ our \ people, \ https://knowledge4policy.ec.europa.eu/publication/communication-com2020562-interval \ outpublication/communication-com2020562-interval \ outpublication/com2020562-interval \ outpublication/communication-com2020562-interval \ outpublication/communication-com2020562-interval \ outpublication/com2020562-interval \ outpublication/com20$
2478	stepping-europe%E2%80%99s-2030-climate-ambition-investing-climate_en, (last access: 10 November 2022), 2020.
2479	
2480	EU: Procedure 2021/0201/COD, COM (2021) 554: Proposal for a REGULATION OF THE EUROPEAN
2481	PARLIAMENT AND OF THE COUNCIL amending Regulations (EU) 2018/841 as regards the scope, simplifying
2482	the compliance rules, setting out the targets of the Member States for 2030 and committing to the collective
2483	achievement of climate neutrality by 2035 in the land use, forestry and agriculture sector, and (EU) 2018/1999 as
2484	regards improvement in monitoring, reporting, tracking of progress and review, https://eur-
2485	lex.europa.eu/procedure/EN/2021_201, 2021a.
2486	
2487	EU: Regulation (EU) 2021/1119 of the European Parliament and of the Council of 30 June 2021 establishing the
2488	framework for achieving climate neutrality and amending Regulations (EC) No 401/2009 and (EU) 2018/1999
2489	('European Climate Law'), https://eur-lex.europa.eu/legal-content/EN/TXT/?uri=CELEX:32021R1119, 2021b.
2490	
2491	EU NIR: Annual European Union greenhouse gas inventory 1990-2019 and inventory report 2021, Submission to the
2492	UNFCCC Secretariat, EEA/PUBL/2021/066, 2021.
2493	
2494	EU NIR: Annual European Union greenhouse gas inventory 1990-2020 and inventory report 2022, Submission to the
2495	UNFCCC Secretariat, EEA/PUBL/2022/023, 2022.
2496	





2497	Federici, S., Tubiello, F. N., Salvatore, M., Jacobs, H., and Schmidhuber, J.: New estimates of CO2 forest emissions
2498	and removals: 1990-2015, Forest Ecol. Manage., 352, 89-98, https://doi.org/10.1016/j.foreco.2015.04.022, 2015.
2499	
2500	Feng, L., Palmer, P. I., Parker, R. J., Deutscher, N. M., Feist, D. G., Kivi, R., Morino, I., and Sussmann, R.: Estimates
2501	of European uptake of CO2 inferred from GOSAT XCO2 retrievals: sensitivity to measurement bias inside and outside
2502	Europe, Atmos. Chem. Phys., 16, 1289–1302, https://doi.org/10.5194/acp-16-1289-2016, 2016.
2503	
2504	Fortems-Cheiney, A., Pison, I., Broquet, G., Dufour, G., Berchet, A., Potier, E., Coman, A., Siour, G., and Costantino,
2505	L.: Variational regional inverse modeling of reactive species emissions with PYVAR-CHIMERE-v2019, Geosci.
2506	Model Dev., 14, 2939–2957, https://doi.org/10.5194/gmd-14-2939-2021, 2021.
2507	
2508	Fortems-Cheiney, A and Broquet, G.: D2.12: Final re-analysis of the national scale CO2 anthropogenic emissions
2509	over 2005-2015, https://projectsworkspace.eu/sites/VERIFY/Deliverables/WP2/VERIFY_D2.12_Final%20re-
2510	analysis%20of%20the%20national%20scale%20CO2%20anthropogenic%20emissions%20over%202005-
2511	2015_v1.pdf, 2021.
2512	
2513	FRA: Global Forest Resources Assessment 2015: How are the world's forest changing?, 2015, Rome, Italy, available
2514	at: http://www.fao.org/3/a-i4793e.pdf (last access: 10 December 2019), 2015.
2515	
2516	Frey, H.C.: Evaluation of an Approximate Analytical Procedure for Calculating Uncertainty in the Greenhouse Gas
2517	Version of the Multi-Scale Motor Vehicle and Equipment Emissions System, Prepared for Office of Transportation
2518	and Air Quality, U.S. Environmental Protection Agency, Ann Arbor, MI, May 30, 2003.
2519	
2520	Friedlingstein, P., O'Sullivan, M., Jones, M. W., Andrew, R. M., Hauck, J., Olsen, A., Peters, G. P., Peters, W.,
2521	Pongratz, J., Sitch, S., Le Quéré, C., Canadell, J. G., Ciais, P., Jackson, R. B., Alin, S., Aragão, L. E. O. C., Arneth,
	A Arora V Datas N. P. Daskar M. Danait Cattin A. Dittig H. C. Dann L. Dultan S. Chandra N. Chavalliar
2522	A., Alora, V., Bates, N. R., Beckel, M., Benon-Cattin, A., Bittig, H. C., Bopp, L., Bunan, S., Chandra, N., Chevanier,
2522 2523	F., Chini, L. P., Evans, W., Florentie, L., Forster, P. M., Gasser, T., Gehlen, M., Gilfillan, D., Gkritzalis, T., Gregor,
2522 2523 2524	 F., Chini, L. P., Evans, W., Florentie, L., Forster, P. M., Gasser, T., Gehlen, M., Gilfillan, D., Gkritzalis, T., Gregor, L., Gruber, N., Harris, I., Hartung, K., Haverd, V., Houghton, R. A., Ilyina, T., Jain, A. K., Joetzjer, E., Kadono, K.,
2522 2523 2524 2525	 F., Chini, L. P., Evans, W., Florentie, L., Forster, P. M., Gasser, T., Gehlen, M., Gilfillan, D., Gkritzalis, T., Gregor, L., Gruber, N., Harris, I., Hartung, K., Haverd, V., Houghton, R. A., Ilyina, T., Jain, A. K., Joetzjer, E., Kadono, K., Kato, E., Kitidis, V., Korsbakken, J. I., Landschützer, P., Lefèvre, N., Lenton, A., Lienert, S., Liu, Z., Lombardozzi,
2522 2523 2524 2525 2526	 K., Arota, V., Bates, N. K., Becker, M., Benote-Catuli, A., Bitug, H. C., Bopp, E., Buitali, S., Chaldra, N., Chevanler, F., Chini, L. P., Evans, W., Florentie, L., Forster, P. M., Gasser, T., Gehlen, M., Gilfillan, D., Gkritzalis, T., Gregor, L., Gruber, N., Harris, I., Hartung, K., Haverd, V., Houghton, R. A., Ilyina, T., Jain, A. K., Joetzjer, E., Kadono, K., Kato, E., Kitidis, V., Korsbakken, J. I., Landschützer, P., Lefèvre, N., Lenton, A., Lienert, S., Liu, Z., Lombardozzi, D., Marland, G., Metzl, N., Munro, D. R., Nabel, J. E. M. S., Nakaoka, SI., Niwa, Y., O'Brien, K., Ono, T., Palmer,





- 2528 Smith, A. J. P., Sutton, A. J., Tanhua, T., Tans, P. P., Tian, H., Tilbrook, B., van der Werf, G., Vuichard, N., Walker,
- 2529 A. P., Wanninkhof, R., Watson, A. J., Willis, D., Wiltshire, A. J., Yuan, W., Yue, X., and Zaehle, S.: Global Carbon
- 2530 Budget 2020, Earth Syst. Sci. Data, 12, 3269–3340, https://doi.org/10.5194/essd-12-3269-2020, 2020.
- 2531
- 2532 Friedlingstein, P., Jones, M. W., O'Sullivan, M., Andrew, R. M., Bakker, D. C. E., Hauck, J., Le Quéré, C., Peters, G.
- 2533 P., Peters, W., Pongratz, J., Sitch, S., Canadell, J. G., Ciais, P., Jackson, R. B., Alin, S. R., Anthoni, P., Bates, N. R.,
- 2534 Becker, M., Bellouin, N., Bopp, L., Chau, T. T. T., Chevallier, F., Chini, L. P., Cronin, M., Currie, K. I., Decharme,
- 2535 B., Djeutchouang, L. M., Dou, X., Evans, W., Feely, R. A., Feng, L., Gasser, T., Gilfillan, D., Gkritzalis, T., Grassi,
- 2536 G., Gregor, L., Gruber, N., Gürses, Ö., Harris, I., Houghton, R. A., Hurtt, G. C., Iida, Y., Ilyina, T., Luijkx, I. T., Jain,
- 2537 A., Jones, S. D., Kato, E., Kennedy, D., Klein Goldewijk, K., Knauer, J., Korsbakken, J. I., Körtzinger, A.,
- 2538 Landschützer, P., Lauvset, S. K., Lefèvre, N., Lienert, S., Liu, J., Marland, G., McGuire, P. C., Melton, J. R., Munro,
- 2539 D. R., Nabel, J. E. M. S., Nakaoka, S.-I., Niwa, Y., Ono, T., Pierrot, D., Poulter, B., Rehder, G., Resplandy, L.,
- 2540 Robertson, E., Rödenbeck, C., Rosan, T. M., Schwinger, J., Schwingshackl, C., Séférian, R., Sutton, A. J., Sweeney,
- 2541 C., Tanhua, T., Tans, P. P., Tian, H., Tilbrook, B., Tubiello, F., van der Werf, G. R., Vuichard, N., Wada, C.,
- 2542 Wanninkhof, R., Watson, A. J., Willis, D., Wiltshire, A. J., Yuan, W., Yue, C., Yue, X., Zaehle, S., and Zeng, J.:
- 2543 Global Carbon Budget 2021, Earth Syst. Sci. Data, 14, 1917–2005, https://doi.org/10.5194/essd-14-1917-2022, 2022.
- 2544
- 2545 Ganzenmüller, R., Bultan, S., Winkler, K., Fuchs, R., Zabel, F., and Pongratz, J.: Land-use change emissions based
- 2546 on high-resolution activity data substantially lower than previously estimated: Environmental Research Letters, 17,
- 2547 64050, DOI 10.1088/1748-9326/ac70d8, 2022.
- 2548
- Gasser, T., Crepin, L., Quilcaille, Y., Houghton, R. A., Ciais, P., and Obersteiner, M.: Historical CO2 emissions from
 land use and land cover change and their uncertainty, Biogeosciences, 17, 4075–4101, https://doi.org/10.5194/bg-174075-2020, 2020.
- 2552
- 2553 Grassi, G., House, J., Kurz, W. A., Cescatti, A., Houghton, R. A., Peters, G. P., Sanz, M. J., Vi nas, R. A., Alkama, .,
- 2554 Arneth, A., Bondeau, A., Dentener, F., Fader, M., Federici, S., Friedlingstein, P., Jain, A. K., Kato, E., Koven, C. D.,
- 2555 Lee, D., Nabel, J. E. M. S., Nassikas, A. A., Perugini, L., Rossi, S., Sitch, S., Viovy, N., Wiltshire, A., and Zaehle, S.:
- 2556 Reconciling global-model estimates and country reporting of anthropogenic forest CO2 sinks, Nat. Clim. Chang., 8,
- 2557 914-920, https://doi.org/10.1038/s41558-018-0283-x, 2018a.
- 2558





- 2559 Grassi, G., Pilli, R., House, J., Federici, S., and Kurz, W. A.: Science-based approach for credible accounting of
- 2560 mitigation in managed forests, Carbon balance Manag., 13, 8, https://doi.org/10.1186/s13021-018-0096-2, 2018b.
- 2561
- 2562 Grassi, G., Cescatti, A., Matthews, R., Duveiller, G., Amia, A., Federici, S., House, J., de Noblet-Ducoudré, N., Pilli,
- 2563 R., and Vizzarri, M.: On the realistic contribution of European forests to reach climate objectives, Carbon balance
- 2564 Manag., 14, 8, https://doi.org/10.1186/s13021-019-0123-y, 2019.
- 2565
- 2566 Grassi, G., Conchedda, G., Federici, S., Abad Viñas, R., Korosuo, A., Melo, J., Rossi, S., Sandker, M., Somogyi, Z.,
- 2567 Vizzarri, M., and Tubiello, F. N.: Carbon fluxes from land 2000–2020: bringing clarity to countries' reporting, Earth
- 2568 Syst. Sci. Data, 14, 4643–4666, https://doi.org/10.5194/essd-14-4643-2022, 2022a.
- 2569
- 2570 Grassi, G., Schwingshackl, C., Gasser, T., Houghton, R. A., Sitch, S., Canadell, J. G., Cescatti, A., Ciais, P., Federici,
- 2571 S., Friedlingstein, P., Kurz, W. A., Sanz Sanchez, M. J., Abad Viñas, R., Alkama, R., Ceccherini, G., Kato, E.,
- 2572 Kennedy, D., Knauer, J., Korosuo, A., McGrath, M. J., Nabel, J., Poulter, B., Rossi, S., Walker, A. P., Yuan, W., Yue,
- 2573 X., and Pongratz, J.: Mapping land-use fluxes for 2001–2020 from global models to national inventories, Earth Syst.
- 2574 Sci. Data Discuss. [preprint], https://doi.org/10.5194/essd-2022-245, in review, 2022b.
- 2575
- 2576 Hansis, E., Davis, S. J., and Pongratz, J.: Relevance of methodological choices for accounting of land use change
- 2577 carbon fluxes, Glob. Biogeochem. Cy., 29, 1230–1246, https://doi.org/10.1002/2014GB004997, 2015.
- 2578
- 2579 Hartung, K., Bastos, A., Chini, L., Ganzenmüller, R., Havermann, F., Hurtt, G. C., Loughran, T., Nabel, J. E. M. S.,
- 2580 Nützel, T., Obermeier, W. A., and Pongratz, J.: Bookkeeping estimates of the net land-use change flux a sensitivity
- study with the CMIP6 land-use dataset, Earth Syst. Dynam., 12, 763–782, https://doi.org/10.5194/esd-12-763-2021,
- **2582** 2021.
- 2583
- Harris, I., Osborn, T.J., Jones, P., and Lister, D.: Version 4 of the CRU TS monthly high-resolution gridded
 multivariate climate dataset, Sci Data 7, 109, https://doi.org/10.1038/s41597-020-0453-3, 2020.
- 2586
- 2587 Hastie, A., Lauerwald, R., Ciais, P., and Regnier, P. : Aquatic carbon fluxes dampen the overall variation of net
- 2588 ecosystem productivity in the Amazon basin: An analysis of the interannual variability in the boundless carbon cycle,
- 2589 Global Change Biology, 25 (6), pp. 2094-2111, DOI: 10.1111/gcb.14620, 2019.




2590	
2591	Haverd, V., Smith, B., Cook, G. D., Briggs, P. R., Nieradzik, L., Roxburgh, S. H., Liedloff, A., Meyer, C. P., and
2592	Canadell, J. G.: A stand-alone tree demography and landscape structure module for Earth system models, Geophysical
2593	Research Letters, 40, 5234-5239, https://doi.org/10.1002/grl.50972, 2013.
2594	
2595	Haverd, V., Smith, B., Nieradzik, L., Briggs, P. R., Woodgate, W., Trudinger, C. M., Canadell, J. G., and Cuntz, M.:
2596	A new version of the CABLE land surface model (Subversion revision r4601) incorporating land use and land cover
2597	change, woody vegetation demography, and a novel optimisation-based approach to plant coordination of
2598	photosynthesis, Geosci. Model Dev., 11, 2995–3026, https://doi.org/10.5194/gmd-11-2995-2018, 2018.
2599	
2600	Houghton, R., Hobbie, J., Melillo, J., Moore, B., Peterson, B., Shaver, G., and Woodwell, G.: Changes in the carbon
2601	content of terrestrial biota and soils between 1860 and 1980: A net release of CO2 to the atmosphere, Ecol. Monogr.,
2602	53, 235–262, https://doi.org/10.2307/1942531, 1983.
2603	
2604	Houghton, R. A.: Revised estimates of the annual net flux of carbon to the atmosphere from changes in land use and
2605	land management 1850-2000, Tellus B, 55, 378-390, https://doi.org/10.3402/tellusb.v55i2.16764, 2003.
2606	
2607	Houghton, R. A., House, J. I., Pongratz, J., van der Werf, G. R., DeFries, R. S., Hansen, M. C., Le Quéré, C., and
2608	Ramankutty, N.: Carbon emissions from land use and land-cover change, Biogeosciences, 9, 5125-5142,
2609	https://doi.org/10.5194/bg-9-5125-2012, 2012.
2610	
2611	Houghton, R. A. and Nassikas, A. A.: Global and regional fluxes of carbon from land use and land cover change
2612	1850-2015, Glob. Biogeochem. Cy., 31, 456-472, https://doi.org/10.1002/2016GB005546, 2017.
2613	
2614	Hurtt, G. C., Chini, L., Sahajpal, R., Frolking, S., Bodirsky, B. L., Calvin, K., Doelman, J. C., Fisk, J., Fujimori, S.,
2615	Klein Goldewijk, K., Hasegawa, T., Havlik, P., Heinimann, A., Humpenöder, F., Jungclaus, J., Kaplan, J. O., Kennedy,
2616	J., Krisztin, T., Lawrence, D., Lawrence, P., Ma, L., Mertz, O., Pongratz, J., Popp, A., Poulter, B., Riahi, K.,
2617	Shevliakova, E., Stehfest, E., Thornton, P., Tubiello, F. N., van Vuuren, D. P., and Zhang, X.: Harmonization of global
2618	land use change and management for the period 850-2100 (LUH2) for CMIP6, Geosci. Model Dev., 13, 5425-5464,
2619	https://doi.org/10.5194/gmd-13-5425-2020, 2020.





2621	IPCC: Good Practice Guidance for Land use, Land use Change and Forestry, Chapter 3, 3.3, https://www.ipcc-
2622	ggip.iges.or.jp/public/gpglulucf/gpglulucf_files/GPG_LULUCF_FULL.pdf, (last access: 10 January 2022), 2003.
2623	
2624	IPCC: Guidelines for National Greenhouse Gas Inventories, Prepared by the National Greenhouse Gas Inventories
2625	Programme. IGES, Japan, https://www.ipcc-nggip.iges.or.jp/public/2006gl/, 2006, (last access: 10 January 2022),
2626	2006.
2627	
2628	IPCC: Refinement to the 2006 IPCC Guidelines for National Greenhouse Gas Inventories, available at:
2629	https://www.ipcc.ch/report/2019-refinement-to-the-2006-ipcc-guidelines-for-national-greenhouse-gas-inventories, the state of the stat
2630	(last access: 10 January 2022), 2019.
2631	
2632	IPCC: Supplement to the 2006 IPCC Guidelines for National Greenhouse Gas Inventories: Wetlands, edited by:
2633	Hiraishi, T., Krug, T., Tanabe, K., Srivastava, N., Baasansuren, J., Fukuda, M., and Troxler, T. G., IPCC, Switzerland,
2634	2014.
2635	
2636	IPCC: Summary for Policymakers. In: Climate Change 2021: The Physical Science Basis.
2637	Contribution of Working Group I to the Sixth Assessment Report of the
2638	Intergovernmental Panel on Climate Change [Masson-Delmotte, V., P. Zhai, A. Pirani,
2639	S.L. Connors, C. Péan, S. Berger, N. Caud, Y. Chen, L. Goldfarb, M.I. Gomis, M. Huang, K.
2640	Leitzell, E. Lonnoy, J.B.R. Matthews, T.K. Maycock, T. Waterfield, O. Yelekçi, R. Yu, and
2641	B. Zhou (eds.)].Cambridge University Press, Cambridge, United Kingdom and New York,
2642	NY, USA, pp. 3-32, doi:10.1017/9781009157896.001, 2021.
2643	
2644	Izaurralde, R. C., Williams, J. R., McGill, W. B., Rosenberg, N.J., and Jakas, M. C. Q.: Simulating soil C dynamics
2645	with EPIC: Model description and testing against long-term data, Ecol. Model. 192, 362-384,
2646	https://doi.org/10.1016/j.ecolmodel.2005.07.010, 2006.
2647	
2648	Janssens-Maenhout, G., Crippa, M., Guizzardi, D., Muntean, M., Schaaf, E., Dentener, F., Bergamaschi, P., Pagliari,
2649	V., Olivier, J. G. J., Peters, J. A. H. W., van Aardenne, J. A., Monni, S., Doering, U., Petrescu, A. M. R., Solazzo, E.,





2650 and Oreggioni, G. D.: EDGAR v4.3.2 Global Atlas of the three major greenhouse gas emissions for the period 1970-2651 2012, Earth Syst. Sci. Data, 11, 959-1002, https://doi.org/10.5194/essd-11-959-2019, 2019. 2652 2653 Jägermeyr, J., Müller, C., Ruane, A.C., Elliott, J., Balkovic, J., Castillo, O., Faye, B., Foster, I., Folberth, C., Franke, 2654 J.A., Fuchs, K., Guarin, J.R., Heinke, J., Hoogenboom, G., Iizumi, T., Jain, A.K., Kelly, D., Khabarov, N., Lange, S., 2655 Lin, T.-S., Liu, W., Mialyk, O., Minoli, S., Moyer, E.J., Okada, M., Phillips, M., Porter, C., Rabin, S.S., Scheer, C., 2656 Schneider, J.M., Schyns, J.F., Skalsky, R., Smerald, A., Stella, T., Stephens, H., Webber, H., Zabel, F., and 2657 Rosenzweig, C.: Climate impacts on global agriculture emerge earlier in new generation of climate and crop models. 2658 Nat. Food, 2, 873-885, https://doi.org/10.1038/s43016-021-00400-y, 2021. 2659 2660 Jenkinson, D. S., Hart, P. B. S., Rayner, J. H., and Parry, L. C.: Modelling the turnover of organic matter in long-term 2661 experiments at Rothamsted, INTECOL Bulletin, 15, 1987. 2662 2663 Jenkinson, D. S., and Rayner, J. H.: The turnover of organic matter in some of the Rothamsted classical experiments, 2664 Soil. Sci., 123, 298-305, https://doi.org/10.1097/00010694-197705000-00005, 1977. 2665 2666 Jonsson, R., Blujdea, V. N., Fiorese, G., Pilli, R., Rinaldi, F., Baranzelli, C., and Camia, A.: Outlook of the European 2667 forest-based sector: forest growth, harvest demand, wood-product markets, and forest carbon dynamics implications, 2668 iForest, 11, 315-328, https://doi.org/10.3832/ifor2636-011, 2018. 2669 Jonsson, R., Rinaldi, F., Pilli, R., Fiorese, G., Hurmekoski, E., Cazzaniga, N., Robert, N., and Camia, A: Boosting the 2670 2671 EU forest-based bioeconomy: Market, climate, and employment impacts, Technological Forecasting and Social 2672 Change, 163, 120478, https://doi.org/10.1016/j.techfore.2020.120478, 2021. 2673 2674 Kanamitsu, M., Ebisuzaki, W., Woollen, J., Yang, S., Hnilo, J. J., Fiorino, M., and Potter, G. L.: NCEP-DOE AMIP-2675 II Reanalysis (R-2). Bulletin of the American Meteorological Society, 83, 1631-1644, https://doi.org/10.1175/BAMS-2676 83-11-1631, 2002. 2677 2678 Klein Goldewijk, K., Beusen, A., Doelman, J., and Stehfest, E.: Anthropogenic land-use estimates for the Holocene; 2679 HYDE 3.2, Earth Syst. Sci. Data, 9, 927-953, https://doi.org/10.5194/essd9-927-2017, 2017a. 2680





2681	Klein Goldewijk, K., Dekker, S. C., and van Zanden, J. L.: Percapita estimations of long-term historical land use and
2682	the consequences for global change research, J. Land Use Sci., 12, 313– 337,
2683	https://doi.org/10.1080/1747423X.2017.1354938, 2017b.
2684	
2685	Koehl, M., Hildebrandt, R., Olschofsky, K., Koehler, R., Roetzer, T., Mette, T., Pretzsch, H., Koethke, M., Dieter,
2686	M., Abiy, M., Makeschin, F., and Kenter, B.: Combating the effects of climatic change on forests by mitigation
2687	strategies, Carbon Balance and Management, 5, 8, https://doi.org/10.1186/1750-0680-5-8, 2010.
2688	
2689	Konovalov, I. B., Berezin, E. V., Ciais, P., Broquet, G., Zhuravlev, R. V., and Janssens-Maenhout, G.: Estimation of
2690	fossil-fuel CO2emissions using satellite measurements of "proxy" species, Atmos. Chem. Phys., 16, 13509-13540,
2691	https://doi.org/10.5194/acp-16-13509-2016, 2016.
2692	
2693	Konovalov, I. B., and Lvova, D. A. : First, fast-track, Re-analysis of the national scale CO2 anthropogenic emissions
2694	over 2005-2015, internal VERIFY report:
2695	$https://projectsworkspace.eu/sites/VERIFY/Deliverables/WP2/VERIFY_D2.10_First,\%20 fast-track,\%20 Re-interval and the second se$
2696	analysis% 20 of% 20 the% 20 national% 20 scale% 20 CO2% 20 anthropogenic% 20 emissions% 20 over% 20 2005-2015.pdf,
2697	(last access: 15 September 2020), 2018.
2698	
2699	Kountouris, P., Gerbig, C., Rödenbeck, C., Karstens, U., Koch, T. F., and Heimann, M.: Technical Note: Atmospheric
2700	CO2 inversions on the mesoscale using data-driven prior uncertainties: methodology and system evaluation, Atmos.
2701	Chem. Phys., 18, 3027–3045, https://doi.org/10.5194/acp-18-3027-2018, 2018a.
2702	
2703	Kountouris, P., Gerbig, C., Rödenbeck, C., Karstens, U., Koch, T. F., and Heimann, M.: Atmospheric CO2 inversions
2704	on the mesoscale using data-driven prior uncertainties: quantification of the European terrestrial CO2 fluxes, Atmos.
2705	Chem. Phys., 18, 3047–3064, https://doi.org/10.5194/acp-18-3047-2018, 2018b.
2706	
2707	Krinner, G., Viovy, N., de Noblet-Ducoudré N., Ogée, J., Polcher, J., Friedlingstein, P., Ciais, P., Sitch, S., and
2708	Prentice, I. C.: A dynamic global vegetation model for studies of the coupled atmosphere-biosphere system, Global
2709	Biogeochemical Cycles, 19, GB1015, doi:10.1029/2003GB002199, 2005.





2711	Kumarathunge, D. P., Medlyn, B. E., Drake, J. E., Tjoelker, M. G., Aspinwall, M. J., Battaglia, M., Cano, F. J., Carter,
2712	K. R., Cavaleri, M. A., Cernusak, L. A., Chambers, J. Q., Crous, K. Y., De Kauwe, M. G., Dillaway, D. N., Dreyer,
2713	E., Ellsworth, D. S., Ghannoum, O., Han, Q., Hikosaka, K., Jensen, A. M., Kelly, J. W. G., Kruger, E. L., Mercado,
2714	L. M., Onoda, Y., Reich, P. B., Rogers, A., Slot, M., Smith, N. G., Tarvainen, L., Tissue, D. T., Togashi, H. F.,
2715	Tribuzy, E. S., Uddling, J., Vårhammar, A., Wallin, G., Warren, J. M. and Way, D. A.: Acclimation and adaptation
2716	components of the temperature dependence of plant photosynthesis at the global scale, New Phytologist, 222, 768-
2717	784, https://doi.org/10.1111/nph.15668, 2019.
2718	
2719	Kurz, W. A., Dymond, C. C., White, T. M., Stinson, G., Shaw, C. H., Rampley, G. J., Smyth, C., Simpson, B. N.,
2720	Neilson, E.T., Trofymow, J. A., Metsaranta, J., and Apps, M. J.: CBMCFS3: a model of carbon dynamics in forestry
2721	and land use change implementing IPCC standards, Ecol. Model., 220, 480-504,
2722	https://doi.org/10.1016/j.ecolmodel.2008.10.018, 2009.
2723	
2724	Lauerwald, R., Laruelle, G. G., Hartmann, J., Ciais, P., and Regnier, P. A. G. :Spatial patterns in CO2 evasion from
2725	the global river network, Global Biogeochemical Cycles, 29, 534-554. https://doi.org/10.1002/2014GB004941015,
2726	2015.
2727	
2728	Lawrence, D. M., Oleson, K. W., Flanner, M. G., Thornton, P. E., Swenson, S. C., Lawrence, P. J., Zeng, X., Yang,
2729	ZL., Levis, S., Sakaguchi, K., Bonan, G. B., and Slater, A. G.: Parameterization Improvements and Functional and
2730	Structural Advances in Version 4 of the Community Land Model, Journal of Advances in Modeling Earth Systems,
2731	3, M03001, DOI 10.1029/2011MS000045, 2011.
2732	
2733	Le Quéré, C., Raupach, M. R., Canadell, J. G., Marland, G., Bopp, L., Ciais, P., Conway, T. J., Doney, S. C., Feely,
2734	R. A., Foster, P., Friedlingstein, P., Gurney, K., Houghton, R. A., House, J. I., Huntingford, C., Levy, P. E., Lomas,
2735	M. R., Majkut, J., Metzl, N., Ometto, J. P., Peters, G. P., Prentice, I. C., Randerson, J. T., Running, S. W., Sarmiento,
2736	J. L., Schuster, U., Sitch, S., Takahashi, T., Viovy, N., van der Werf, G. R., and Woodward, F. I.: Trends in the sources
2737	and sinks of carbon dioxide, Nat. Geosci., 2, 831-836, https://doi.org/10.1038/ngeo689, 2009.
2738	
2739	Liski, J., Palosuo, T., Peltoniemi, M., and Sievänen, R.: Carbon and decomposition model Yasso for forest soils, Ecol.
2740	Model., 189, 168-182, https://doi.org/10.1016/J.ECOLMODEL.2005.03.005, 2005.

2741





2742	Liu, J., Baskaran, L., Bowman, K., Schimel, D., Bloom, A. A., Parazoo, N. C., Oda, T., Carroll, D., Menemenlis, D.,
2743	Joiner, J., Commane, R., Daube, B., Gatti, L. V., McKain, K., Miller, J., Stephens, B. B., Sweeney, C., and Wofsy,
2744	S.: Carbon Monitoring System Flux Net Biosphere Exchange 2020 (CMS-Flux NBE 2020), Earth Syst. Sci. Data, 13,
2745	299-330, https://doi.org/10.5194/essd-13-299-2021, 2021.
2746	
2747	Lugato, E., Panagos, P., Bampa, F., Jones, A., Montanarella, L.: A new baseline of organic carbon stock in European
2748	agricultural soils using a modeling approach, Glob. Change Biol., 20, 313-326, https://doi.org/10.1111/gcb.12292,
2749	2014.
2750	
2751	Lurton, T., Balkanski, Y., Bastrikov, V., Bekki, S., Bopp, L., Braconnot, P., Brockmann, P., Cadule, P., Contoux, C.,
2752	Cozic, A., Cugnet, D., Dufresne, JL., Éthé, C., Foujols, MA., Ghattas, J., Hauglustaine, D., Hu, RM., Kageyama,
2753	M., Khodri, M., Lebas, N., Levavasseur, G., Marchand, M., Ottlé, C., Peylin, P., Sima, A., Szopa, S., Thiéblemont,
2754	R., Vuichard, N., and Boucher, O.: Implementation of the CMIP6 Forcing Data in the IPSL-CM6A-LR Model. Journal
2755	of Advances in Modeling Earth Systems, 12(4), e2019MS001940, https://doi.org/10.1029/2019MS001940, 2020.
2756	
2757	Luyssaert, S., Abril, G., Andres, R., Bastviken, D., Bellassen, V., Bergamaschi, P., Bousquet, P., Chevallier, F., Ciais,
2758	P., Corazza, M., Dechow, R., Erb, KH., Etiope, G., Fortems-Cheiney, A., Grassi, G., Hartmann, J., Jung, M.,
2759	Lathière, J., Lohila, A., Mayorga, E., Moosdorf, N., Njakou, D. S., Otto, J., Papale, D., Peters, W., Peylin, P.,
2760	Raymond, P., Rödenbeck, C., Saarnio, S., Schulze, ED., Szopa, S., Thompson, R., Verkerk, P. J., Vuichard, N.,
2761	Wang, R., Wattenbach, M., and Zaehle, S.: The European land and inland water CO2, CO, CH4 and N2O balance
2762	between 2001 and 2005, Biogeosciences, 9, 3357-3380, https://doi.org/10.5194/bg-9-3357-2012, 2012.
2763	
2764	Luyssaert, S., Marie, G., Valade, A., Chen, Y. Y., Njakou Djomo, S., Ryder, J., Otto, J., Naudts, K., Lansø, A. S.,
2765	Ghattas, J., and McGrath, M. J.: Trade-offs in using European forests to meet climate objectives, Nature, 562, 259-
2766	262, https://doi.org/10.1038/s41586-018-0577-1, 2018.
2767	
2768	Mason Earles, J., Yeh, S. and Skog, K: Timing of carbon emissions from global forest clearance, Nature Clim Change,
2769	2, 682-685, https://doi.org/10.1038/nclimate1535, 2012.
2770	
2771	McGrath, M. J., Petrescu, A. M. R., Peylin, P., Andrew, R. M., Matthews, B., Dentener, F., Balkovič, J., Bastrikov,

2772 V., Becker, M., Broquet, G., Ciais, P., Fortems, A., Ganzenmüller, R., Grassi, G., Harris, I., Jones, M., Knauer, J.,





Kuhnert, M., Monteil, G., Munassar, S., Palmer, P. I., Peters, G. P., Qiu, C., Schelhaas, MJ., Tarasova, O., Vizzarri,
M., Winkler, K., Balsamo, G., Berchet, A., Briggs, P., Brockmann, P., Chevallier, F., Conchedda, G., Crippa, M.,
Dellaert, S., Denier van der Gon, H. A. C., Filipek, S., Friedlingstein, P., Fuchs, R., Gauss, M., Gerbig, C., Guizzardi,
D., Günther, D., Houghton, R. A., Janssens-Maenhout, G., Lauerwald, R., Lerink, B., Luijkx, I. T., Moulas, G.,
Muntean, M., Nabuurs, GJ., Paquirissamy, A., Perugini, L., Peters, W., Pilli, R., Pongratz, J., Regnier, P., Scholze,
M., Serengil, Y., Smith, P., Solazzo, E., Thompson, R. L., Tubiello, F. N., Vesala, T. and Walther, S.: Data for the
consolidated European synthesis of CO2 emissions and removals for EU27 and UK: 1990-2020,
https://doi.org/10.5281/zenodo.7365863, 2022.
Menut, L., Bessagnet, B., Khvorostyanov, D., Beekmann, M., Blond, N., Colette, A., Coll, I., Curci, G., Foret, G.,
Hodzic, A., Mailler, S., Meleux, F., Monge, JL., Pison, I., Siour, G., Turquety, S., Valari, M., Vautard, R., and
Vivanco, M. G.: CHIMERE 2013: a model for regional atmospheric composition modeling, Geosci. Model Dev., 6,
981-1028, https://doi.org/10.5194/gmd-6-981-2013, 2013.
Messager, M. L., Lehner, B., Grill, G., Nedeva, I. and Schmitt, O.: Estimating the volume and age of water stored in
global lakes using a geo-statistical approach, Nat. Commun., 7, 13603, doi:10.1038/ncomms13603, 2016.
Monteil, G., Broquet, G., Scholze, M., Lang, M., Karstens, U., Gerbig, C., Koch, FT., Smith, N. E., Thompson, R.
L., Luijkx, I. T., White, E., Meesters, A., Ciais, P., Ganesan, A. L., Manning, A., Mischurow, M., Peters, W., Peylin,
P., Tarniewicz, J., Rigby, M., Rödenbeck, C., Vermeulen, A., and Walton, E. M.: The regional European atmospheric
transport inversion comparison, EUROCOM: first results on European-wide terrestrial carbon fluxes for the period
2006–2015, Atmos. Chem. Phys., 20, 12063–12091, https://doi.org/10.5194/acp-20-12063-2020, 2020.
Monteil, G., and Scholze, M.: Regional CO2 inversions with LUMIA, the Lund University Modular Inversion
Algorithm, v1.0, Geoscientific Model Development, 14, 3383-3406, https://doi.org/10.5194/gmd-14-3383-2021,
2021.
Mueller, N., Gerber, J., Johnston, M., Ray, D. K., Ramankutty, N., and Foley, J. A.: Closing yield gaps through nutrient
and water management, Nature, 490, 254-257, https://doi.org/10.1038/nature11420, 2012.





2803	Muñoz-Sabater, J.: ERA5-Land hourly data from 1981 to present, Copernicus Climate Change Service (C3S) Climate
2804	Data Store (CDS), DOI 10.24381/cds.e2161bac, (last access: 1 May 2021), 2019.
2805	
2806	Muñoz-Sabater, J., Dutra, E., Agustí-Panareda, A., Albergel, C., Arduini, G., Balsamo, G., Boussetta, S., Choulga,
2807	M., Harrigan, S., Hersbach, H., Martens, B., Miralles, D. G., Piles, M., Rodríguez-Fernández, N. J., Zsoter, E.,
2808	Buontempo, C., and Thépaut, JN.: ERA5-Land: A state-of-the-art global reanalysis dataset for land applications,
2809	Earth Syst. Sci. Data, 13, 4349-4383, https://doi.org/10.5194/essd-13-4349-2021, 2021.
2810	
2811	Nabuurs, G, Lindner, M, Verkerk, H, Gunia, K, Deda, P, Michalak, R, and Grassi, G.: First signs of carbon sink
2812	saturation in European forest biomass, Nature Climate Change 3, 792-796, https://doi.org/10.1038/nclimate1853,
2813	2013.
2814	
2815	Nabuurs, G. J., Delacote, P., Ellison, D., Hanewinkel, M., Hetemäki, L., Lindner, M., and Ollikainen, M.: By 2050
2816	the mitigation effects of EU forests could nearly double through climate smart forestry, Forests, 8, 484,
2817	https://doi.org/10.3390/f8120484, 2017.
2818	
2819	Nabuurs, G. J., Arets, E. J. M. M., and Schelhaas, M. J.: Understanding the implications of the EU-LULUCF regulation
2820	for the wood supply from EU forests to the EU, Carbon Balance Manag., 13, 18, https://doi.org/10.1186/s13021-018-
2821	0107-3, 2018.
2822	
2823	Naegler, T.: Reconciliation of excess 14C-constrained global CO2 piston velocity estimates, Tellus B, 61, 372-384,
2824	https://doi.org/10.1111/j.1600-0889.2008.00408.x, 2009.
2825	
2826	Naudts, K., Chen, Y., McGrath, M., Ryder, J., Valade, A., Otto, J., and Luyssaert, S.: Europe's forest management
2827	did not mitigate climate warming, Science, 351, 597-600, https://doi.org/10.1126/science.aad7270, 2016.
2828	
2829	Niwa, Y., Fujii, Y., Sawa, Y., Iida, Y., Ito, A., Satoh, M., Imasu, R., Tsuboi, K., Matsueda, H., and Saigusa, N.: A
2830	4D-Var inversion system based on the icosahedral grid model (NICAM-TM 4D-Var v1.0) - Part 2: Optimization
2831	scheme and identical twin experiment of atmospheric CO2 inversion, Geosci. Model Dev., 10, 2201-2219,
2832	https://doi.org/10.5194/gmd-10-2201-2017, 2017.
2833	





2834 Oleson, K.: Technical Description of the Community Land Model (CLM). NCAR Technical Note. TN-478+STR.

- 2835 10.5065/D6RR1W7M, 2010.
- 2836
- 2837 Petrescu, A. M. R., Peters, G. P., Janssens-Maenhout, G., Ciais, P., Tubiello, F. N., Grassi, G., Nabuurs, G.-J., Leip,
- 2838 A., Carmona-Garcia, G., Winiwarter, W., Höglund-Isaksson, L., Günther, D., Solazzo, E., Kiesow, A., Bastos, A.,
- 2839 Pongratz, J., Nabel, J. E. M. S., Conchedda, G., Pilli, R., Andrew, R. M., Schelhaas, M.-J., and Dolman, A. J.:
- 2840 European anthropogenic AFOLU greenhouse gas emissions: a review and benchmark data, Earth Syst. Sci. Data, 12,
- 2841 961–1001, https://doi.org/10.5194/essd-12-961-2020, 2020.
- 2842
- 2843 Petrescu, A. M. R., McGrath, M. J., Andrew, R. M., Peylin, P., Peters, G. P., Ciais, P., Broquet, G., Tubiello, F. N.,
- 2844 Gerbig, C., Pongratz, J., Janssens-Maenhout, G., Grassi, G., Nabuurs, G.-J., Regnier, P., Lauerwald, R., Kuhnert, M.,
- 2845 Balkovič, J., Schelhaas, M.-J., Denier van der Gon, H. A. C., Solazzo, E., Qiu, C., Pilli, R., Konovalov, I. B.,
- 2846 Houghton, R. A., Günther, D., Perugini, L., Crippa, M., Ganzenmüller, R., Luijkx, I. T., Smith, P., Munassar, S.,
- 2847 Thompson, R. L., Conchedda, G., Monteil, G., Scholze, M., Karstens, U., Brockmann, P., and Dolman, A. J.: The
- 2848 consolidated European synthesis of CO2 emissions and removals for the European Union and United Kingdom: 1990–
- 2849 2018, Earth Syst. Sci. Data, 13, 2363–2406, https://doi.org/10.5194/essd-13-2363-2021, 2021b.
- 2850
- 2851 Pilli, R., Grassi, G., Kurz, W. A., Moris, J. V., and Viñas, R. A.: Modelling forest carbon stock changes as affected
- 2852 by harvest and natural disturbances II. EU-level analysis including land use changes, Carbon Balance and
- 2853 Management, 11, 20, https://doi.org/10.1186/s13021-016-0059-4, 2016.
- 2854
- Pilli, R., Grassi, G., Kurz, W. A., Fiorese, G., and Cescatti, A.: The European forest sector: past and future carbon
 budget and fluxes under different management scenarios, Biogeosciences, 14, 2387–2405, https://doi.org/10.5194/bg14-2387-2017, 2017.
- 2858
- Pilli, R., Alkama, R., Cescatti, A., Kurz, W. A., and Grassi, G.: The European forest carbon budget under future
 climate conditions and current management practices, Biogeosciences, 19, 3263–3284, https://doi.org/10.5194/bg-193263-2022, 2022.
- 2862
- 2863 Pinty B., Janssens-Maenhout, G., Dowell, M., Zunker, H., Brunhes, T., Ciais, P., Dee, D., Denier van der Gon, H.,
- 2864 Dolman, H., Drinkwater, M., Engelen, R., Heimann, M., Holmlund, K., Husband, R., Kentarchos, A., Meijer, Y.,





- Palmer, P., and Scholze, M.: An Operational Anthropogenic CO₂ Emissions Monitoring & Verification Support
 capacity Baseline Requirements, Model Components and Functional Architecture, European Commission Joint
 Research Centre, EUR 28736 EN, doi: 10.2760/39384, 2017.
- 2868
- 2869 Pisso, I., Sollum, E., Grythe, H., Kristiansen, N. I., Cassiani, M., Eckhardt, S., Arnold, D., Morton, D., Thompson, R.
- 2870 L., Groot Zwaaftink, C. D., Evangeliou, N., Sodemann, H., Haimberger, L., Henne, S., Brunner, D., Burkhart, J. F.,
- 2871 Fouilloux, A., Brioude, J., Philipp, A., Seibert, P., and Stohl, A.: The Lagrangian particle dispersion model
- 2872 FLEXPART version 10.4, Geosci. Model Dev., 12, 4955–4997, https://doi.org/10.5194/gmd-12-4955-2019, 2019.
- 2873
- 2874 Polcher, J., McAvaney, B., Viterbo, P., Gaertner, M.-A., Hahmann, A., Mahfouf, J.-F., Noilhan, J., Phillips, T.,
- Pitman, A.J., Schlosser, C.A., Schulz, J.-P., Timbal, B., Verseghy D., and Xue, Y.: A proposal for a general interface
 between land-surface schemes and general circulation models, Global and Planetary Change, 19, 263-278,
 https://doi.org/10.1016/S0921-8181(98)00052-6, 1998.
- 2878
- Pongratz, J., Reick, C. H., Houghton, R. A., and House, J. I.: Terminology as a key uncertainty in net land use and
 land cover change carbon flux estimates, Earth Syst. Dynam., 5, 177–195, https://doi.org/10.5194/esd-5-177-2014,
 2014.

2882

- Pongratz, J., Schwingshackl, C., Bultan, S., Obermeier, W., Havermann, F., and Guo, S.: Land Use Effects on Climate:
 Current State, Recent Progress, and Emerging Topics. Curr Clim Change Rep, 7, 99–120,
 https://doi.org/10.1007/s40641-021-00178-y, 2021.
- 2886
- Pongratz, J., Reick, C., Raddatz, T., and Claussen, M.: A reconstruction of global agricultural areas and land cover
 for the last millennium, Global Biogeochemical Cycles, 22, https://doi.org/10.1029/2007GB003153, 2008.
- 2889
- Ramankutty, N., and Foley, J. A.: Estimating historical changes in global land cover: Croplands from 1700 to 1992,
 Global biogeochemical cycles, 13, 997-1027, https://doi.org/10.1029/1999GB900046, 1999.
- 2892
- Raymond, P. A., Hartmann, J., Lauerwald, R., Sobek, S., McDonald, C., Hoover, M., and Guth, P. : Global carbon dioxide emissions from inland waters, Nature, 503, 355–359, https://doi.org/10.1038/nature12760, 2013.
- 2896 RECAPP2: https://www.globalcarbonproject.org/Reccap/index.htm, last access: 22 November 2022.
- 2897





2898	Regnier, P., Friedlingstein, P., Ciais, P., Mackenzie, F. T., Gruber, N., Janssens, I. A., Laruelle, G. G., Lauerwald, R.,
2899	Luyssaert, S., Andersson, A. J., Arndt, S., Arnosti, C., Borges, A. V., Dale, A. W., Gallego-Sala, A., Goddéris, Y.,
2900	Goossens, N., Hartmann, J., Heinze, C., Ilyina, T., Joos, F., LaRowe, D. E., Leifeld, J., Meysman, F. J. R., Munhoven,
2901	G., Raymond, P. A., Spahni, R., Suntharalingam, P., and Thullner, M.: Anthropogenic perturbation of the carbon
2902	fluxes from land to ocean, Nature Geosci, 6, 597-607, https://doi.org/10.1038/ngeo1830, 2013.
2903	
2904	Reichstein, M., Bahn, M., Ciais, P., Frank, D., Mahecha, M. D., Seneviratne, S. I., Zscheischler, J., Beer, C.,
2905	Buchmann, N., Frank, D. C., Papale, D., Rammig, A., Smith, P., Thonicke, K., van der Velde, M., Vicca, S., Walz,
2906	A., and Wattenbach, M.: Climate extremes and the carbon cycle, Nature, 500, 287-95, doi: 10.1038/nature12350,
2907	2013.
2908	
2909	Resplandy, L., Keeling, R.F., Rödenbeck, C. Stephens, B. B., Khatiwala, S., Rodgers, K. B., Long, M. C., Bopp, L.,
2910	and Tans, P. P.: Revision of global carbon fluxes based on a reassessment of oceanic and riverine carbon transport,
2911	Nature Geosci, 11, 504–509, https://doi.org/10.1038/s41561-018-0151-3, 2018.
2912	
2913	Rödenbeck, C.: Estimating CO2 sources and sinks from atmospheric mixing ratio measurements using a global
2914	inversion of atmospheric transport, Tech. Rep. 6, Max Planck Institute for Biogeochemistry, Jena, Germany, 2005.
2915	
2916	Rödenbeck, C., Gerbig, C., Trusilova, K., and Heimann, M.: A two-step scheme for high-resolution regional
2917	atmospheric trace gas inversions based on independent models, Atmos. Chem. Phys., 9, 5331-5342,
2918	https://doi.org/10.5194/acp-9-5331-2009, 2009.
2919	
2920	Rödenbeck, C., Bakker, D. C., Metzl, N., Olsen, A., Sabine, C., Cassar, N., Reum, F., Keeling, R. F. and Heimann,
2921	M.: Interannual sea-air CO2 flux variability from an observation-driven ocean mixed-layer scheme, Biogeosciences,
2922	11, 4599-4613, https://doi.org/10.5194/bg-11-4599-2014, 2014.
2923	
2924	Sallnäs, O.: A matrix model of the Swedish forest, Studia Forestalia Suecica, 183, 1-23,
2925	https://pub.epsilon.slu.se/4514/, 1990.
2926	
2927	Scarlat, N, Martinov, M, and Dallemand, J.F.: Assessment of the availability of agricultural crop residues in the
2928	European Union: potential and limitations for bioenergy use, Waste Manag, 10, 1889-97, doi:
2929	10.1016/j.wasman.2010.04.016, 2010.
2930	
2931	Scharnweber, T., Smiljanic, M., Cruz-García, R., Manthey, M., and Wilmking, M.: Tree growth at the end of the 21st
2932	century - the extreme years 2018/19 as template for future growth conditions, Environ. Res. Lett., 15, 074022,
2933	https://doi.org/10.1088/1748-9326/ab865d, 2020.
2934	





2935	Schelhaas, MJ., Nabuurs, GJ., Verkerk, P.J., Hengeveld, G., Packalen, T., Sallnäs, O., Pilli, R., Grassi, G., Forsell,
2936	N., Frank, S., Gusti, M., and Havlik, P.: Forest Resource Projection Tools at the European Level. In: Barreiro, S.,
2937	Schelhaas, MJ., McRoberts, R.E., Kändler, G. (Eds.), Forest Inventory-based Projection Systems for Wood and
2938	Biomass Availability, Springer International Publishing, Cham, pp. 49-68, 2017.
2939	
2940	Schelhaas, M. J., Hengeveld, G. M., Filipek, S., König, L., Lerink, B., Staristsky, I., de Jong, A., Sikkema, R., and
2941	Nabuurs, G. J.: Documentation of the EFISCEN Space model, in prep.
2942	
2943	Seidl, R., Schelhaas, M. J., Rammer, W. and Verkerk, P. J.: Increasing forest disturbances in Europe and their impact
2944	on carbon storage, Nature Clim Change, 4, 806-810, https://doi.org/10.1038/nclimate2318, 2014.
2945	
2946	Silva, J. P., Toland, J., Jones, W., Eldrige, J., Thorpe, E., O'Hara, E.: LIFE and Europe's grasslands: Restoring a
2947	forgotten habitat, report by the European Commission,
2948	$https://ec.europa.eu/environment/archives/life/publications/lifepublications/lifefocus/documents/grassland.pdf, \ (last the second se$
2949	access: 10 November 2022), 2008.
2950	
2951	Simmonds, P., Palmer, P. I., Rigby, M., McCulloch, A., O'Doherty, S. G., and Manning, A. J.: Tracers for evaluating
2952	computational models of atmospheric transport and dispersion at regional to global scales, Atm. Env., 246, 118074,
2953	doi:10.1016/j.atmosenv.2020.118074, 2021.
2954	
2955	Simpson, D., Benedictow, A., Berge, H., Bergström, R., Emberson, L. D., Fagerli, H., Flechard, C. R., Hayman, G.
2956	D., Gauss, M., Jonson, J. E., Jenkin, M. E., Nyíri, A., Richter, C., Semeena, V. S., Tsyro, S., Tuovinen, JP.,
2957	Valdebenito, Á., and Wind, P.: The EMEP MSC-W chemical transport model - technical description, Atmos. Chem.
2958	Phys., 12, 7825-7865, doi:10.5194/acp-12-7825-2012, 2012.
2959	
2960	Simpson, D., Bergström, R., Imhof, H., and Wind, P.: Updates to the emep/msc-w model, 2016-2017. In
2961	Transboundary particulate matter, photo-oxidants, acidifying and eutrophying components. EMEP Status Report
2962	1/2017. The Norwegian Meteorological Institute, Oslo, Norway, 2017.
2963	
2964	Simpson, D., Bergström, R., Tsyro, S., and Wind, P.: Updates to the EMEP MSC-W model, 2018- 2019. In
2965	Transboundary particulate matter, photo-oxidants, acidifying and eutrophying components. EMEP Status Report
2966	1/2019. The Norwegian Meteorological Institute, Oslo, Norway, 2019.
2967	
2968	Simpson, D., Gonzalez Fernandez, I. A., Segers, A., Tsyro, S., Valdebenito, A., and Wind, P.: Updates to the EMEP
2969	MSC-W model, 2021-2022. In Transboundary particulate matter, photo-oxidants, acidifying and eutrophying
2970	components. EMEP Status Report 1/2022. The Norwegian Meteorological Institute, Oslo, Norway, 2022.
2971	





2972 De Smet, P. A. M., and Hettelingh, J.-P.: Intercomparison of Current European Land Use/Land Cover Databases, 2973 Status Report 2001 Coordination Center for Effects, RIVM Report 259101010, Bilthoven, Netherlands, pp. 41-52, 2974 2001. 2975 2976 Smith, J. U., Bradbury, N. J., and Addiscott, T.M.: SUNDIAL: A PC-based system for simulating nitrogen dynamics 2977 in arable land, Agron J, 88, 38-43, https://doi.org/10.2134/agronj1996.00021962008800010008x, 1996. 2978 2979 Smith, J. U., Gottschalk, P., Bellarby, J., Chapman, S., Lilly, A., Towers, W., Bell, J., Coleman, K., Nayak, D. R., 2980 Richards, M. I., Hillier, J., Flynn, H. C., Wattenbach, M., Aitkenhead, M., Yeluripurti, J. B., Farmer, J., Milne, R., 2981 Thomson, A., Evans, C., Whitmore, A. P., Falloon, P. and Smith, P.: Estimating changes in national soil carbon stocks 2982 using ECOSSE - a new model that includes upland organic soils. Part I. Model description and uncertainty in national 2983 scale simulations of Scotland, Climate Research, 45, 179-192, doi: 10.3354/cr00899, 2010a. 2984 2985 Smith, J.U., Gottschalk, P., Bellarby, J., Chapman, S., Lilly, A., Towers, W., Bell, J., Coleman, K., Nayak, D. R., 2986 Richards, M. I., Hillier, J., Flynn, H. C., Wattenbach, M., Aitkenhead, M., Yeluripurti, J. B., Farmer, J., Milne, R., 2987 Thomson, A., Evans, C., Whitmore, A.P., Falloon, P. and Smith, P.: Estimating changes in national soil carbon stocks 2988 using ECOSSE - a new model that includes upland organic soils. Part II Application in Scotland, Climate Research, 2989 45, 193-205, doi: 10.3354/cr00902, 2010b. 2990 2991 Smith, B., Wärlind, D., Arneth, A., Hickler, T., Leadley, P., Siltberg, J., and Zaehle, S.: Implications of incorporating 2992 N cycling and N limitations on primary production in an individual-based dynamic vegetation model, Biogeosciences, 2993 11, 2027–2054, https://doi.org/10.5194/bg-11-2027-2014, 2014. 2994 2995 Solazzo, E., Crippa, M., Guizzardi, D., Muntean, M., Choulga, M., and Janssens-Maenhout, G.: Uncertainties in the 2996 Emissions Database for Global Atmospheric Research (EDGAR) emission inventory of greenhouse gases, Atmos. 2997 Chem. Phys., 21, 5655-5683, https://doi.org/10.5194/acp-21-5655-2021, 2021. 2998 2999 Thompson, R. L., Broquet, G., Gerbig, C., Koch, T., Lang, M., Monteil, G., Munassar, S., Nickless, A., Scholze, M., 3000 Ramonet, M., Karstens, U., van Schaik, E., Wu, Z. and Rödenbeck, C.: Changes in net ecosystem exchange over 3001 Europe during the 2018 drought based on atmospheric observations, Phil. Trans. R. Soc. B, 375, 20190512, 3002 http://dx.doi.org/10.1098/rstb.2019.0512, 2020. 3003 3004 Toreti, A., Belward, A., Perez-Dominguez, I., Naumann, G., Luterbacher, J., Cronie, O., Lorenzo Seguini, L., 3005 Manfron, G., Lopez-Lozano, R., Baruth, B., van den Berg, M., Dentener, F., Ceglar, A., Chatzopoulos, T., and 3006 Zampieri, M.: The exceptional 2018 European water seesaw calls for action on adaptation, Earth's Future, 7, 652-663, 3007 https://doi.org/10.1029/2019EF001170, 2019.





3009	UK NIR: UK Greenhouse Gas Inventory, 1990 to 2020, Annual Report for Submission under the Framework
3010	Convention on Climate Change, 978-0-9933975-8-5, 2022.
3011	
3012	UNFCCC: Kyoto Climte Change Decision, available at: https://unfccc.int/process-and-meetings/conferences/past-
3013	conferences/kyoto-climate-change-conference-december-1997/decisions-kyoto-climate-change-conference-
3014	december-1997, (last access: 5 October 2020), 1997.
3015	
3016	UNFCCC: Decision 24/CP.19 Revision of the UNFCCC reporting guidelines on annual inventories for Parties
3017	included in Annex I to the Convention, FCCC/CP/2013/10/Add.3, 2014.
3018	
3019	UNFCCC: NGHGI 2021 NIR reports: National Inventory Submissions 2021, available at: https://unfccc.int/ghg-
3020	inventories-annex-i-parties/2021, (last access: 01 January 2022), 2022a.
3021	
3022	UNFCCC: NGHGI 2021 CRFs, available at: https://unfccc.int/ghg-inventories-annex-i-parties/2021, (last access 01
3023	March 2022), 2022b.
3024	
3025	VERIFY: http://verify.lsce.ipsl.fr/, last access: 21 November 2022.
3026 3027 3028	VERIFY Synthesis Plots: http://webportals.ipsl.jussieu.fr/VERIFY/FactSheets/, last access: 21 November 2022.
3029	van der Laan-Luijkx, I. T., van der Velde, I. R., van der Veen, E., Tsuruta, A., Stanislawska, K., Babenhauserheide,
3030	A., Zhang, H. F., Liu, Y., He, W., Chen, H., Masarie, K. A., Krol, M. C., and Peters, W.: The CarbonTracker Data
3031	Assimilation Shell (CTDAS) v1.0: implementation and global carbon balance 2001–2015, Geosci. Model Dev., 10,
3032	2785-2800, https://doi.org/10.5194/gmd-10-2785-2017, 2017.
3033	
3034	Verkerk, P. J., Schelhaas, MJ., Immonen, V., Hengeveld, G., Kiljunen, J., Lindner, M., Nabuurs, GJ., Suominen,
3035	T., and Zudin, S.: Manual for the European Forest Information Scenario model (EFISCEN 4.1), EFI Technical Report
3036	99, European Forest Institute, 49 pp., 2016.
3037	
3038	Viovy, N.: Interannuality and CO2 sensitivity of the SECHIBA-BGC coupled SVAT-BGC model, Physics and
3039	Chemistry of The Earth, 21, 489-497, https://doi.org/10.1016/S0079-1946(97)81147-0, 1996.
3040	
3041	Vizzarri, M., Pilli, R., Korosuo, A., Blujdea, V. N. B., Rossi, S., Fiorese, G., Abad-Vinas, R., Colditz, R. R., and
3042	Grassi, G.: Setting the forest reference levels in the European Union: overview and challenges, Carbon Balance
3043	Manage, 16, 23, https://doi.org/10.1186/s13021-021-00185-4, 2021.
3044	





3045	Wang, YP., and Leuning, R.: A two-leaf model for canopy conductance, photosynthesis and partitioning of available
3046	energy I: Model description and comparison with a multi-layered model, Agricultural and Forest Meteorology, 91,
3047	89-111, https://doi.org/10.1016/S0168-1923(98)00061-6, 1998.
3048	
3049	Wang, Y. P., Law, R. M., and Pak, B.: A global model of carbon, nitrogen and phosphorus cycles for the terrestrial
3050	biosphere, Biogeosciences, 7, 2261-2282, https://doi.org/10.5194/bg-7-2261-2010, 2010.
3051	
3052	Wanninkhof, R.: Relationship between wind speed and gas exchange over the ocean revisited: Gas exchange and wind
3053	speed over the ocean, Limnol. OceanogrMeth., 12, 351-362, https://doi.org/10.4319/lom.2014.12.351, 2014.
3054	
3055	Williams, J. R.: The Erosion-Productivity Impact Calculator (EPIC) Model: A Case History, Philos. Trans. R. Soc. B
3056	Biol. Sci. 329, 421–428, https://doi.org/10.1098/rstb.1990.0184, 1990.
3057	
3058	Winkler, K., Fuchs, R., Rounsevell, M. D. A., and Herold, M.: HILDA+ Global Land Use Change between 1960 and
3059	2019. PANGAEA, https://doi.org/10.1594/PANGAEA.921846, 2020.
3060	
3061	Winkler, K., Fuchs, R., Rounsevell, M. and Herold, M.: Global land use changes are four times greater than previously
3062	estimated, Nat Commun, 12, 2501, https://doi.org/10.1038/s41467-021-22702-2, 2021.
3063	
3064	WMO: United in Science Report, available at: https://public.wmo.int/en/our-mandate/climate/wmo-statement-state-
3065	of-global-climate, (last access: January 2022), 2021.
3066	
3067	Yvon-Durocher, G., Caffrey, J., Cescatti, A., Dossena, M., del Giorgio, P., Gasol, J. M., Montoya, J. M., Pumpanen,
3068	J., Staehr, P. A., Trimmer, M., Woodward, G., and Allen, A. P.: Reconciling the temperature dependence of respiration
3069	across timescales and ecosystem types, Nature, 487, 472-476, https://doi.org/10.1038/nature11205, 2012.
3070	
3071	Zhang, B., Tian, H., Lu, C., Dangal, S. R. S., Yang, J., and Pan, S.: Global manure nitrogen production and application
3072	in cropland during 1860-2014: a 5 arcmin gridded global dataset for Earth system modeling, Earth Syst. Sci. Data, 9,
3073	667-678, https://doi.org/10.5194/essd-9-667-2017, 2017.
3074	
3075	Zscheischler, J., Mahecha, M. D., Avitabile, V., Calle, L., Carvalhais, N., Ciais, P., Gans, F., Gruber, N., Hartmann,
3076	J., Herold, M., Ichii, K., Jung, M., Landschützer, P., Laruelle, G. G., Lauerwald, R., Papale, D., Peylin, P., Poulter,
3077	B., Ray, D., Regnier, P., Rödenbeck, C., Roman-Cuesta, R. M., Schwalm, C., Tramontana, G., Tyukavina, A.,
3078	Valentini, R., van der Werf, G., West, T. O., Wolf, J. E., and Reichstein, M.: Reviews and syntheses: An empirical
3079	spatiotemporal description of the global surface-atmosphere carbon fluxes: opportunities and data limitations,
3080	Biogeosciences, 14, 3685–3703, https://doi.org/10.5194/bg-14-3685-2017, 2017.
3081	