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1 **LegacyClimate 1.0: A dataset of pollen-based climate**
2 **reconstructions from 2594 Northern Hemisphere sites**
3 **covering the late Quaternary**

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53 **Abstract.** Here we describe the LegacyClimate 1.0, a dataset of the reconstruction of mean July
54 temperature (T_{July}), mean annual temperature (T_{ann}), and annual precipitation (P_{ann}) from 2594 fossil
55 pollen records from the Northern Hemisphere spanning the entire Holocene with some records reaching
56 back to the Last Glacial. Two reconstruction methods, the Modern Analogue Technique (MAT) and
57 Weighted-Averaging Partial-Least Squares regression (WA-PLS) reveal similar results regarding spatial
58 and temporal patterns. To reduce the impact of precipitation on temperature reconstruction and vice
59 versa, we also provide reconstructions using tailored modern pollen data limiting the range of the
60 corresponding other climate variable. We assess the reliability of the reconstructions using information
61 from the spatial distributions of the root-mean squared error of prediction and reconstruction significance
62 tests. The dataset is beneficial for climate proxy synthesis studies and to evaluate the output of climate
63 models and thus help to improve the models themselves. We provide our compilation of reconstructed
64 T_{July} , T_{ann} , and P_{ann} as open-access datasets at PANGAEA
65 (<https://doi.pangaea.de/10.1594/PANGAEA.930512>; Herzschuh et al., 2021). R code for the
66 reconstructions is provided at Zenodo (<https://doi.org/10.5281/zenodo.5910989>; Herzschuh et al., 2022),
67 including harmonized open-access modern and fossil datasets used for the reconstructions, so that
68 customized reconstructions can be easily established.

69

70 **1 Introduction**

71 The evaluation of climate model outputs using climate data is essential for model improvements (Eyring
72 et al., 2019). However, the period for which observations are available is only of limited use to validate
73 simulations because it is short and characterized by strong changes in the climate driver. Climate proxy
74 data derived from natural archives are therefore of great value. The extratropical Northern Hemisphere
75 is of particular interest because it is known for complex spatial and temporal temperature and
76 precipitation patterns.

77 Previous proxy-based climate inferences have contributed to major debates about Holocene climate
78 change. For example, while simulations indicate a gradual warming of the Holocene, temperature proxy
79 data syntheses rather support a mid-Holocene optimum which resulted in the “Holocene conundrum”



80 debate (Liu et al., 2014). Qualitative proxy-based inferences indicate that the mid-Holocene in the
81 Northern Hemisphere mid-latitudes was rather dry and warm compared with present-day in agreement
82 with modeling outputs (Routson et al., 2019). Also, quantitative precipitation reconstructions from
83 Eastern and Central Asia unveiled the complex monsoon-westerlies interactions (Chen et al., 2019;
84 Herzschuh et al., 2019).

85 Fossil pollen records are well-established in their use as a palaeoecological and palaeoclimatological
86 proxy and of great value as indicators of past environmental and climatic change for many decades.
87 Considerable efforts have been made to establish regional, continental and even global data repositories
88 like the North American Pollen Database (<http://www.ncdc.noaa.gov/paleo/napd.html>), the European
89 Pollen Database (<http://www.europeanpollendatabase.net>) and the Neotoma Paleoecology Database
90 (<https://www.neotomadb.org>; Williams et al., 2018). Regarding the prevalence of pollen archives across
91 multiple environmental settings such as lakes, wetlands, or marine sediments, fossil pollen records are
92 widely used to quantitatively reconstruct past vegetation and climate variables (Birks, 2019; Chevalier
93 et al., 2020). Pollen data are the only land-derived proxy data that have sufficient temporal and spatial
94 coverage to allow for high-resolution climate model evaluation of the late Quaternary period. A number
95 of methods have been proposed for making pollen-based climate reconstructions (Chevalier et al., 2020):
96 among them, classification approaches like the Modern Analogue Technique (MAT) or regression
97 approaches like Weighted-Averaging Partial-Least Squares regression (WA-PLS) are most commonly
98 used.

99 For temperature reconstruction time-series, several broad-scale syntheses exist; however, either they
100 originate from different proxies (Kaufman et al., 2020a and 2020b) or are restricted to certain continents
101 or regions (Mauri et al., 2015; Marsicek et al., 2018; Routson et al., 2019). Temperature reconstructions
102 from the large extratropical Asia are mostly lacking. Precipitation syntheses are available from Europe
103 (Mauri et al., 2015), North America (Whitmore et al., 2005) and China and Mongolia (Herzschuh et al.,
104 2019) but, hitherto, no global or hemispheric syntheses of quantitative precipitation changes are
105 available for the Holocene.

106 In a recent effort, pollen records available in the Neotoma Paleoecology Database (Williams et al.,
107 2018) and additional records from China and Siberia (Cao et al., 2013 and 2020) were synthesized and
108 taxonomically harmonized (Herzschuh et al., submitted). Furthermore, all chronologies of these records
109 were recently revised using a Bayesian approach that allows for the inference of temporal uncertainties



110 (Li et al., 2022). Here we present the pollen-based reconstruction of mean July temperature (T_{July}), mean
111 annual temperature (T_{ann}) and annual precipitation (P_{ann}) from these 2594 records from the Northern
112 Hemisphere using WA-PLS and MAT.

113

114 **2 Methods**

115 **2.1 Input data**

116 The objective of this study is to create a dataset of quantitative reconstructions of T_{July} , T_{ann} and P_{ann}
117 spanning the Holocene from a set of fossil pollen records. We used fossil data from the Neotoma
118 Paleocology Database (Williams et al., 2018; <https://www.neotomadb.org>; downloaded in July 2020),
119 a dataset from Eastern and Central Asia (Cao et al., 2013; Herzschuh et al., 2019) and a dataset from
120 Northern Asia (Cao et al., 2020). The harmonized dataset is stored on PANGAEA (LegacyPollen 1.0)
121 and presented in Herzschuh et al. (submitted). Ages were taken from age-depth models presented in Li
122 et al. (2022), who recently provided a set of harmonized chronologies under the “LegacyAge 1.0”
123 framework, and applied to our fossil pollen synthesis. A modern pollen training dataset comprised of
124 15,379 sites includes datasets from Eurasia (EMPD1, Davis et al. 2013; EMPD2, Davis et al. 2020;
125 Herzschuh et al., 2019; Tarasov et al., 2011) and North America (Whitmore et al., 2005). In order to
126 reduce inconsistencies in pollen identification, the modern and fossil pollen datasets were taxonomically
127 harmonized: major tree and shrub pollen were merged to genus level and most of the herbaceous taxa
128 (except the most common ones such as *Artemisia*, *Thalictrum* or *Rumex*) to family level. We excluded
129 aquatic pollen (with the exception of *Cyperaceae*), spores from ferns and fungi, as well as algae and
130 calculated pollen percentages on the basis of the total number of terrestrial pollen grains. The site
131 specific T_{ann} , T_{July} , P_{ann} were derived from WorldClim 2 (spatial resolution of 1 km,
132 <https://www.worldclim.org>, Fick and Hijmans, 2017) by extracting the climate data at the location of the
133 modern sample sites using the *raster* package in R (version 3.5-11, Hijmans et al., 2021; R Core Team,
134 2020).

135 We compiled the fossil data into four sub-continental datasets for Eastern North America ($<105^{\circ}\text{W}$;
136 Williams et al., 2000), Western North America, Europe and Asia. For consistency with the amount of
137 taxa in the North American training dataset, the fossil datasets were reduced to the 70 most common



138 taxa on the respective sub-continent, according to Hill's N2 diversity index (i.e., the effective number
139 of occurrences of a species in the dataset; Hill, 1973).

140

141 **2.2 Reconstruction methods**

142 Our reconstruction approach included MAT (Overpeck et al., 1985) and WA-PLS (ter Braak and Juggins,
143 1993) by applying the *MAT* and *WAPLS* functions from the *rioja* package (version 0.9-21, Juggins, 2019)
144 for R (R Core Team, 2020) on our Northern Hemispheric fossil pollen synthesis. For each fossil location,
145 we calculated the geographic distance between each modern sampling site and the fossil pollen record
146 using the *rdist.earth* function from the *fields* R-package (version 10.3, Nychka et al., 2020) and selected
147 a unique calibration set from modern sites within a 2000 km radius. For the reconstruction with MAT,
148 we used the original pollen percentages of the selected fossil pollen taxa, looking for 7 analogues
149 between the pollen data and the selected calibration dataset. The dissimilarity between the fossil
150 samples and the modern pollen assemblages was determined by squared-chord distance metrics
151 (Simpson, 2012; Cao et al., 2014). For the reconstruction with WA-PLS, we used the square-root
152 transformed pollen percentages in a leave-one-out cross-validation approach (Cao et al., 2014). In
153 addition to the classic WA-PLS reconstruction, we provide WA-PLS_tailored. To reconstruct T_{July} we
154 “tailored” our modern training dataset with respect to the P_{ann} range. For this purpose, we identified the
155 range of the reconstructed P_{ann} and extended it by 25% to both ends of the modern P_{ann} range in order
156 to reduce the influence of P_{ann} on T_{ann} and T_{July} reconstruction due to co-variation. We applied the same
157 method to the reconstruction of P_{ann} . T_{ann} and T_{July} were tailored by P_{ann} ; P_{ann} was tailored by T_{July} and,
158 additionally, by T_{ann} (illustrated for an example in Appendix Fig. 1). A statistical significance test (Telford
159 and Birks, 2011) was performed for the reconstruction by using the *randomTF* function in the *palaeoSig*
160 R-package (version 2.0-3, Telford, 2019). The reconstructed climate parameters were tested as single
161 variables, as well as with partialling out the respective other variable. We applied a Canonical Correlation
162 Analysis (CCA) to the modern training dataset in order to infer the explained variance in the modern
163 dataset by using the *cca* function in the *vegan* R-package (version 2.5-7, Oksanen et al., 2020). The
164 ratio between constrained (λ_1) and unconstrained (λ_2) explained variance was determined for all modern
165 training datasets.



166 **3 Dataset description LegacyClimate 1.0: input data, reconstructions and reconstruction model**
 167 **statistics**

168 LegacyClimate 1.0 provides pollen-based reconstructions and sample-specific reconstruction errors of
 169 T_{ann} , T_{July} and P_{ann} for 2594 fossil pollen records (i.e., a total of 146,067 single pollen samples) from
 170 three reconstruction methods (WA-PLS, WA-PLS_tailored, MAT). Furthermore, we provide the method-
 171 specific model metadata and statistics for each record and each climate variable (Table 1). To ease
 172 data handling, the dataset files are separated into Western North America, Eastern North America,
 173 Europe and Asia.

174 **Table 1.** Structure and content of the LegacyClimate 1.0 data with details about the information
 175 contained in the input, datasets, in the reconstructions and the reconstruction model statistics.

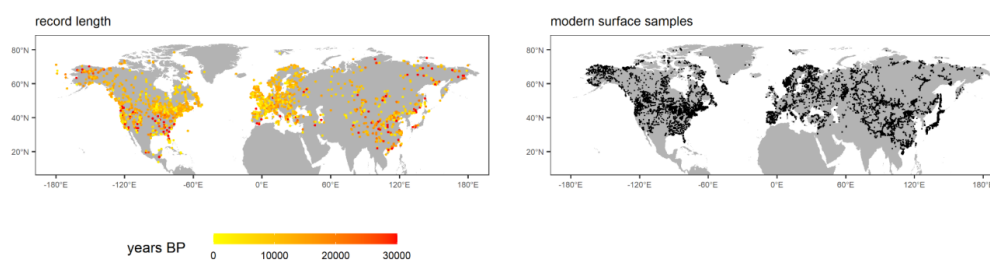
Datasets	Content
Input datasets	Modern pollen dataset of 15,379 sites
	Modern dataset of T_{ann} , T_{July} , P_{ann}
	Fossil pollen data (LegacyPollen 1.0) for 2594 sites with a total of 146,067 samples
LegacyClimate 1.0: Climate reconstructions	Reconstructions and sample-specific reconstruction errors of T_{ann} , T_{July} and P_{ann} for 2594 sites using MAT, WA-PLS and WA-PLS_tailored
LegacyClimate 1.0: Reconstruction model statistics	Site information (Event label, Source, ID, Site name, Longitude, Latitude)
	Modern pollen dataset information (number of modern analogues, range of climate variables)
	Model statistics for each site for MAT, WA-PLS, WA-PLS_tailored (including r^2 observed vs. predicted, RMSEP, no. of WA-PLS components)



176 4 Dataset assessment

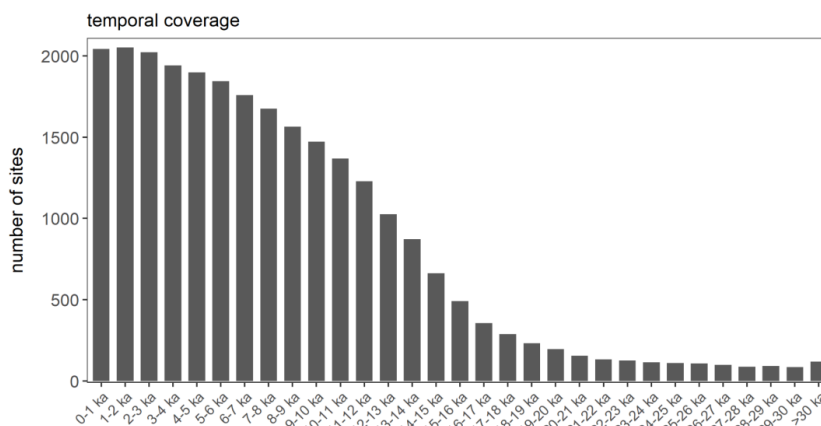
177 4.1 Spatial and temporal coverage of LegacyClimate 1.0

178 In total, we provide reconstructions for 2594 fossil pollen records, among them 670 records from Eastern
179 North America, 361 records from Western North America, 1075 records from Europe and 488 Asian
180 records (Fig. 1). The temporal coverage of the records is rather uneven: 119 and 289 records cover the
181 periods before 30,000 years (Fig. 2) and the Last Glacial Maximum, respectively. A total of 1229, 1845,
182 2052 records are available for 12-11 ka, 6-5 ka BP and 2-1 ka BP, respectively.



183

184 **Figure 1.** left: map indicating the spatial distribution and record lengths covered by the LegacyPollen
185 1.0 dataset (Herzschuh et al., submitted) for which climate reconstructions are provided in
186 LegacyClimate 1.0 with a total of 2594 records; right: spatial distribution of modern pollen dataset used
187 for reconstruction with a total of 15,379 sites.



188

189 **Figure 2.** Number of records that cover certain millennia of the last 30 ka.

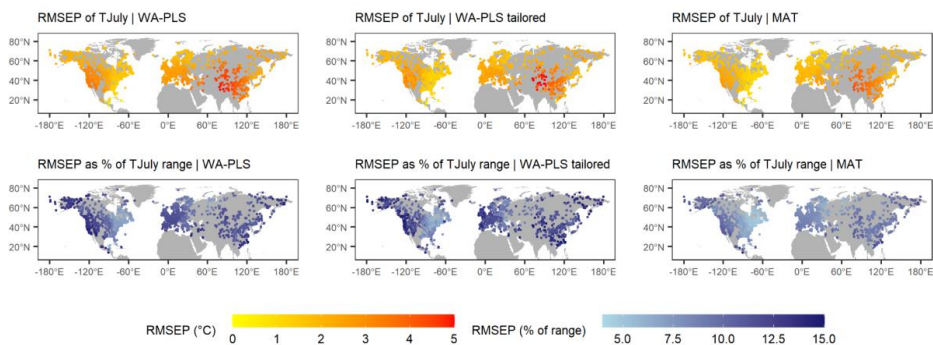


190 **4.2 Prediction errors of LegacyClimate 1.0**

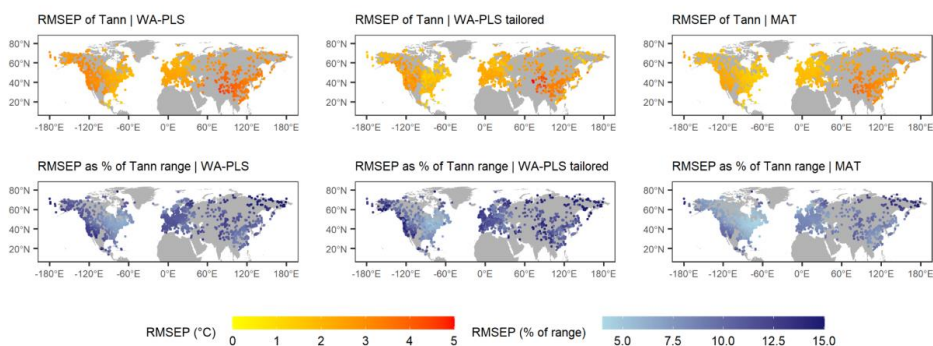
191 The mean RMSEPs and their standard deviations for T_{ann} are $1.98 \pm 0.52^\circ\text{C}$ (MAT), $2.61 \pm 0.53^\circ\text{C}$ (WA-
192 PLS) and $2.24 \pm 0.61^\circ\text{C}$ (WA-PLS_tailored) and mean RMSEPs as a percentage of modern T_{ann} range
193 are $7.68 \pm 1.93\%$ (MAT), $10.09 \pm 2.05\%$ (WA-PLS) and $10.26 \pm 2.79\%$ (WA-PLS_tailored). The largest
194 mean RMSEP values are located in Central Asia in Kazakhstan, Mongolia and the north-western parts
195 of the Tibetan Plateau and are consistent across all three reconstruction methods. Other areas with
196 large mean RMSEP values are located in Western North America, Southern and Central Europe and
197 south-east Asia. The smallest RMSEPs can be found along the east coast of North America. Relative
198 to the modern temperature range, the RMSEP from this region also reveals the lowest fraction. In
199 general, MAT has the lowest mean error fraction relative to the modern temperature range of all three
200 methods.

201 The mean RMSEPs of T_{July} are $1.90 \pm 0.63^\circ\text{C}$ (MAT), $2.50 \pm 0.73^\circ\text{C}$ (WA-PLS) and $2.21 \pm 0.75^\circ\text{C}$ (WA-
202 PLS_tailored) and mean percentages of T_{July} range are $8.11 \pm 1.64\%$ (MAT), $10.71 \pm 1.94\%$ (WA-PLS)
203 and $10.70 \pm 2.60\%$ (WA-PLS_tailored). Thus, they are slightly smaller than those of T_{ann} but slightly larger
204 as a percentage of the range. The spatial patterns, however, are largely similar to those of T_{ann} .

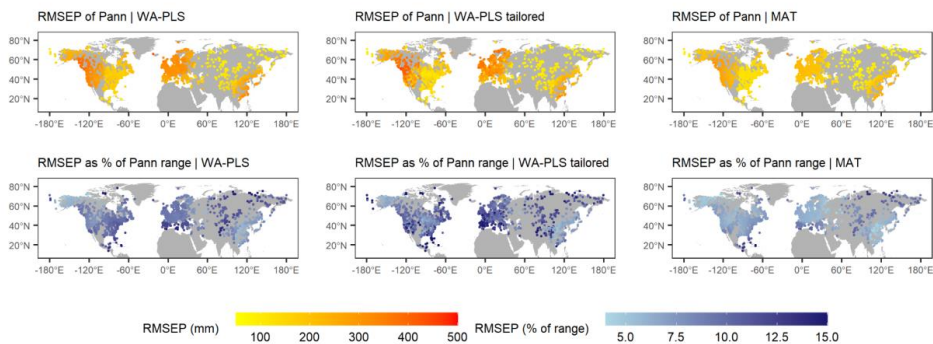
205 The mean RMSEPs of P_{ann} are 176.38 ± 51.40 mm (MAT), 244.48 ± 75.84 mm (WA-PLS) and
206 232.71 ± 98.57 mm (WA-PLS_tailored) and mean percentages of P_{ann} range are $6.78 \pm 1.48\%$ (MAT),
207 $9.27 \pm 1.70\%$ (WA-PLS) and $10.26 \pm 2.67\%$ (WA-PLS_tailored). High RMSEPs are found for Western
208 North America, Europe and along the coastline of south-east Asia, while the lowest RMSEP values are
209 found for Central Asia. A clear division in RMSEPs are found on the North American continent: while
210 the western part of North America (with the exception of Alaska) has a rather high RMSEP, the eastern
211 part of North America has a smaller RMSEP. This pattern is found for all three methods (Fig. 3).



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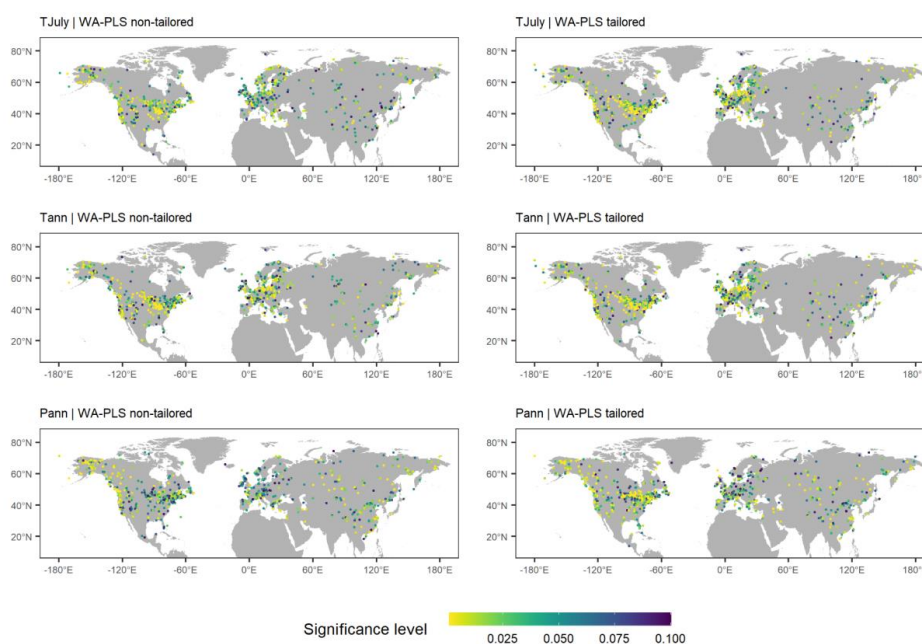
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215 **Figure 3.** Spatial distribution of root mean squared error of prediction (RMSEP) as inferred from leave-
216 one out cross-validation presented as absolute values and as a percentage of the range of mean July
217 temperature (T_{July}), mean annual temperature (T_{ann}), mean annual precipitation (P_{ann}) in the modern
218 pollen data used for reconstruction for the three methods applied (weighted-averaging partial least



219 squares (WA-PLS), WA-PLS using a training set from within a limited climate range (WA-PLS_tailored)
220 and modern analogue technique (MAT)).

221 A significance test ($p < 0.1$) according to Telford and Birks (2011) for the whole reconstructed time
222 period was run for each record and for the reconstructions with WA-PLS and WA-PLS_tailored (Fig. 4;
223 Table 2). The T_{July} reconstruction is significant for 30.9% (WA-PLS) and 35.2% (WA-PLS_tailored) when
224 included as a single variable in the significance test. Partialling out precipitation as a conditional variable
225 causes an increase in the amount of significant records to 35.5% for WA-PLS, but a decrease for WA-
226 PLS_tailored to 33.6% of all records. For T_{ann} , 32.8% (WA-PLS) and 36.1% (WA-PLS_tailored) of all
227 records pass the significance test when tested as a single variable. When partialling out precipitation,
228 the amount of significant records decreases for both WA-PLS and WA-PLS_tailored. 32.1% (WA-PLS)
229 and 33.4% (WA-PLS_tailored) of all records pass the significance test when testing P_{ann} as a single
230 variable. In contrast to the significance tests for T_{ann} , partialling out the mean July temperature as a
231 conditional variable increases the number of significant records for both WA-PLS and WA-PLS_tailored.





233 **Figure 4.** Maps showing mean July temperature (T_{July}), mean annual temperature (T_{ann}), mean annual
234 precipitation (P_{ann}) records that passed the reconstruction significance test ($p < 0.1$). Color indicates the
235 significance level.

236 **Table 2.** Percentage of records that pass the reconstruction significance test ($p < 0.1$) sensu Telford and
237 Birks (2011).

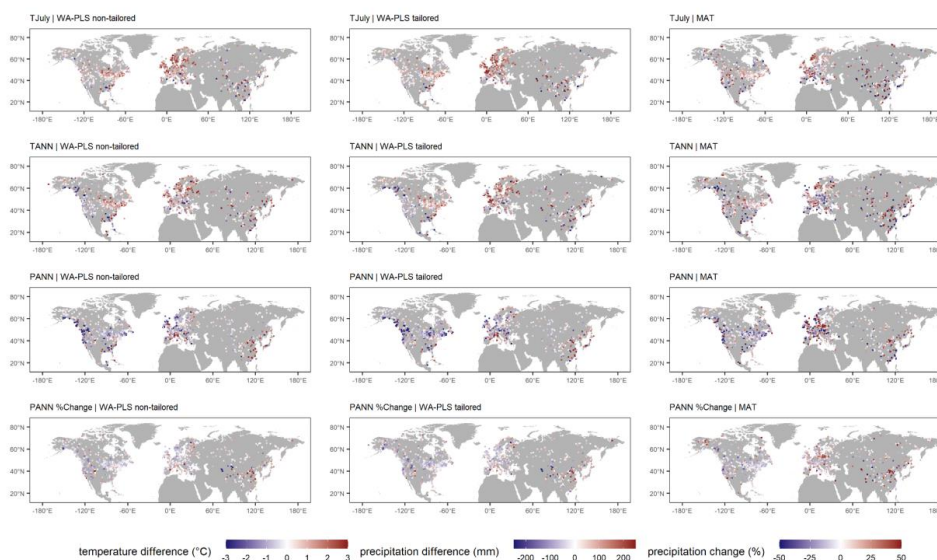
	WA-PLS	WA-PLS_tailored
T_{July}	30.9%	35.2%
T_{July} partialling out P_{ann}	35.5%	33.6%
T_{ann}	32.8%	36.1%
T_{ann} partialling out P_{ann}	32.6%	34.1%
P_{ann}	32.1%	33.4%
P_{ann} partialling out T_{July}	34.3%	36.5%

238

239 4.3 Assessment of major temporal patterns of LegacyClimate 1.0

240 For analyzing the temporal variation, we calculated means of all three climate variables for the time
241 periods between 6.5 and 5.5 ka BP and between 1.5 and 0.5 ka BP and subtracted those means from
242 every record in order to evaluate the changes between the reconstructed mid-Holocene conditions and
243 those of modern times. Differences between these time periods reveal warmer and drier conditions
244 especially in Eastern North America but also in Central and Northern Europe. The overall patterns are
245 in good agreement for all three methods but show differences on a regional scale, especially when
246 comparing the reconstructions with WA-PLS and MAT. For T_{July} , the reconstruction with MAT shows
247 greater temperature differences in Western North America and south-east Asia. Compared to the
248 reconstruction with WA-PLS, there is a reduced cooling in Eastern Europe and a warming in the Western
249 Mediterranean region and along the south-eastern Asian coastline. Comparing the reconstructions of
250 T_{ann} , more gradual patterns are seen in the reconstruction with WA-PLS: Western North America reveals
251 a mid-Holocene warming, while Eastern North America shows a cooling. In Europe records that report
252 a cooling are more concentrated in the northern and western parts of the continent. In the reconstruction
253 with MAT, Eastern North America is divided into a reported cooling in the northern part and a warming

254 in the southern part. In Western North America, there is a mixture of locations with a warming and a
255 cooling since the mid-Holocene. In Europe, only France and Southern Scandinavia show a cooling; in
256 Central and parts of Southern Europe, a warming can be found in the reconstructions. For large areas
257 in North America and Europe, the reconstructions with WA-PLS suggest an increase in precipitation
258 since 6 ka BP. A shift to drier conditions can be found along the south-eastern coastline in North America,
259 in the Mediterranean Region and especially in south-east Asia. The reconstruction with MAT reveals a
260 gradient from increasing precipitation in south-western Europe to decreasing precipitation in north-
261 eastern Europe. In contrast to the reconstructions with WA-PLS, records along the south-eastern Asian
262 coastline suggest an increase in precipitation with MAT rather than a decrease (Fig. 5).



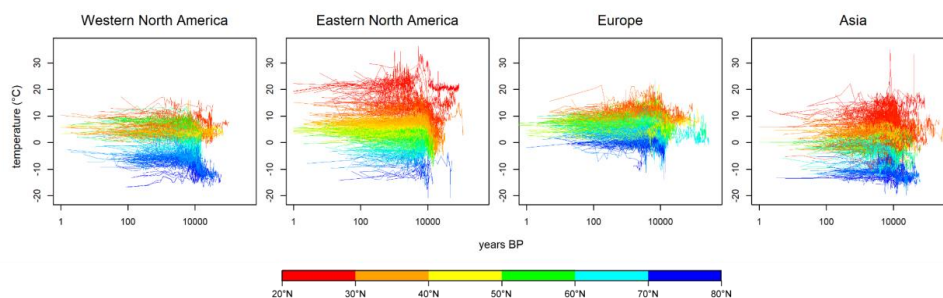
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264 **Figure 5.** Difference from 6 ka to 1 ka for mean July temperature (T_{July}), mean annual temperature (T_{ann}),
265 mean annual precipitation (P_{ann}) and $P_{ann}\%$ as reconstructed from weighted-averaging partial least
266 squares (WA-PLS), WA-PLS using a training set from within a limited climate range (WA-PLS_tailored)
267 and modern analogue technique (MAT).

268 Time-series of absolute T_{ann} reconstructions reveal temporal as well as latitudinal spatial variation on
269 the single continents. Eastern North America and Asia show the most variation in the low latitudes. It is
270 also Eastern North America which shows the most pronounced latitudinal gradient. In Western North
271 America, the most variation takes place in the high latitudes, while the variation is concentrated to the

272 mid-latitudes in Europe. Especially in North America, the warming since the last deglaciation and the
273 beginning of the Holocene is well shown in the temporal variation of the time-series (Fig. 6).

274



275

276 **Figure 6.** Time-series of absolute mean annual temperature (T_{ann}) reconstruction for each
277 (sub-)continent. Colors denote the latitude of record origin. Note logarithmic x-axis.

278

279 4.4 Assessment of consistency among reconstruction methods

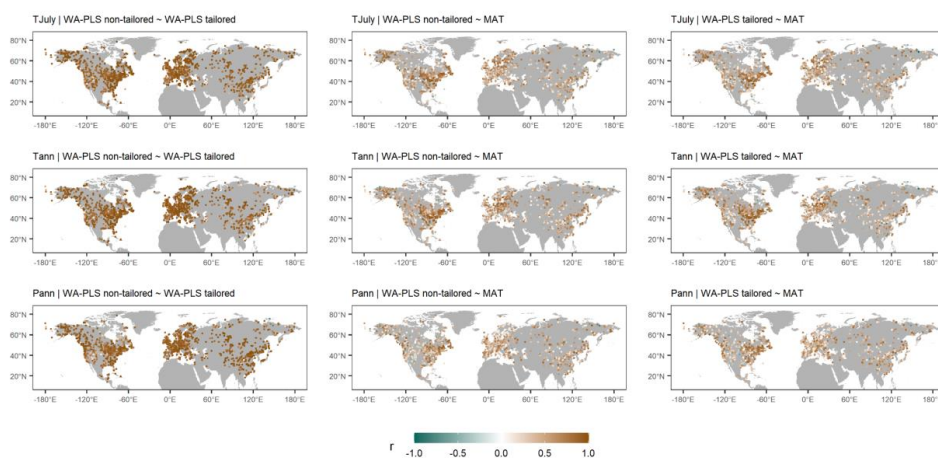
280 Reconstructions with MAT are, in general, in good agreement with those derived from the WA-PLS.

281 Comparing MAT with WA-PLS, 37.3% (T_{July}), 38.9% (T_{ann}) and 30.4% (P_{ann}) of all records have a positive

282 correlation of $r \geq 0.6$. Strong positive correlations ($r \geq 0.9$) can mainly be identified in Eastern North

283 America, while weak correlation can be found for large areas in central North America and most of

284 Europe (Fig. 7).



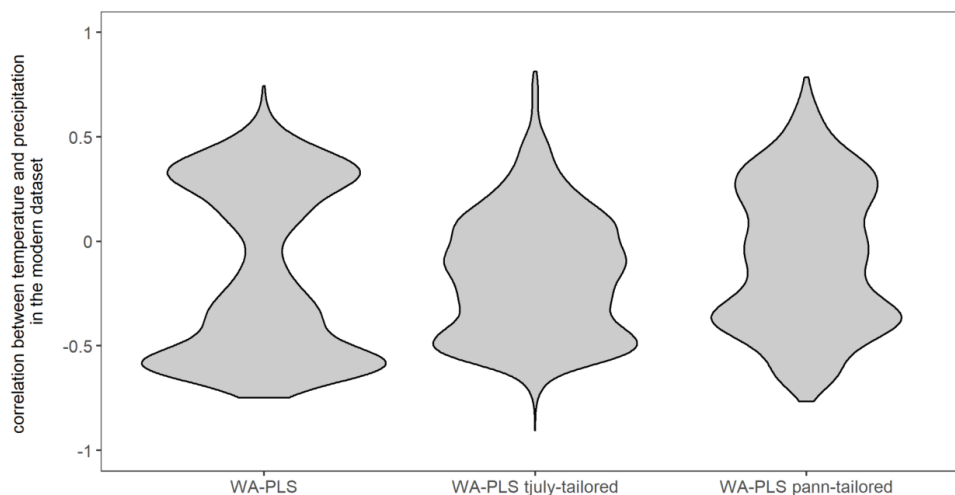
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286 **Figure 7.** Correlation between time-series of the 3 different reconstruction methods used - weighted-
287 averaging partial least squares using a global training set (WA-PLS), WA-PLS using a training set with
288 a limited modern climate range (WA-PLS_tailored) and the modern analogue technique (MAT) for the
289 three climate variables of mean July temperature (T_{July}), mean annual temperature (T_{ann}) and mean
290 annual precipitation (P_{ann})

291

292 WA-PLS_tailored used a reduced modern training dataset (illustrated for an example in Appendix Fig.
293 1). The tailoring successfully reduced the co-variation of temperature and precipitation in the modern
294 dataset as indicated by the distribution of the correlation coefficient in Fig. 8. Nevertheless, the obtained
295 reconstructions are largely consistent between WA-PLS and WA-PLS-tailored: a correlation of $r \geq 0.9$
296 is found for 59.2% of all records for T_{July} , 60.7% for T_{ann} and 56.5% for P_{ann} .

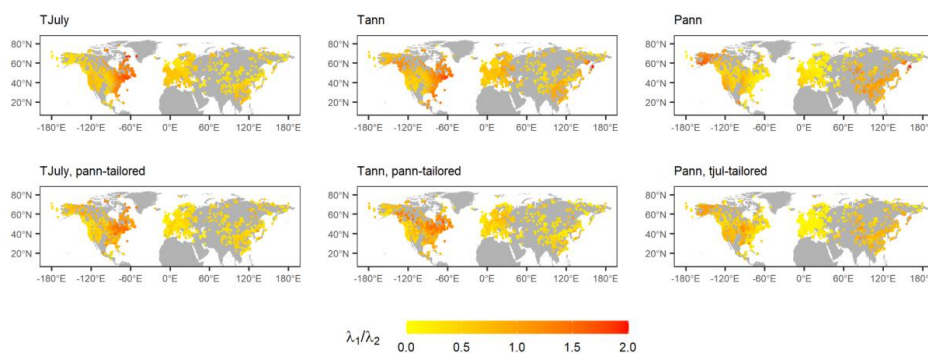


297

298 **Figure 8.** Violin plot of the correlation coefficients between T_{July} and P_{ann} in the 15,379 training datasets
299 used for the reconstructions. Left: used for WA-PLS reconstructions; middle: WA-PLS tjuly-tailored
300 (used for the reconstruction of P_{ann}); WA-PLS pann-tailored (used for the reconstruction of T_{July}).

301 A CCA was performed to infer the ratio between constrained and unconstrained explained variance for
302 all modern training datasets (λ_1/λ_2) for the modern datasets used for WA-PLS and WA-PLS_tailored.
303 Modern datasets used for WA-PLS constrained by T_{July} reveal a concentration of high ratios in Eastern
304 North America while low ratios can be found in Central Asia. While the spatial pattern of λ_1/λ_2 constrained

305 by T_{ann} is similar, the ratios are slightly higher for T_{ann} than for T_{July} . Reconstructions for P_{ann} show low
306 ratios in Europe and Eastern North America. Areas with high ratios are concentrated in Alaska and East
307 Asia.



308

309 **Figure 9.** Maps showing λ_1/λ_2 , representing the ratio of explained variance of first axis (constrained) vs.
310 second (unconstrained) axis as revealed by applying a CCA to all modern training datasets that were
311 used for the reconstructions. Constraining variable as well as tailoring of the dataset (see methods) is
312 indicated in the map captions.

313

314 5 Discussion

315 5.1 Impact of the fossil pollen data source on LegacyClimate 1.0 quality

316 LegacyClimate 1.0 contains reconstructions of climate variables from fossil pollen data derived from
317 open-access data repositories. The fossil records were derived from multiple natural archives, most
318 commonly, assemblages from continuous lacustrine and peat accumulations (Herzschuh et al.,
319 submitted). Different sizes of lakes and peat areas result in varying sizes of pollen source areas and
320 thus the spatial representativeness of a record, as small lakes and peatlands are considered to provide
321 information about the (extra-)local scale, while pollen assemblages from large lakes are considered as
322 a regional signal (Jackson, 1990; Sugita, 1993). However, such signals might be impacted by
323 taphonomy of the record, for example pollen from azonal riverine vegetation might be over-represented
324 in fluvially impacted pollen records.

325 Our dataset is based on taxonomically harmonized modern and fossil pollen datasets using a restricted
326 number of taxa (i.e., the most common 70 taxa on each (sub-)continent). Such an approach guarantees



327 that all records are handled consistently. Although losing taxonomic information when merging taxa
328 together into a higher taxonomic level, it also increases the possibility of matching climate analogues in
329 the modern and the fossil datasets. However, one needs to keep in mind that species with different
330 ecological requirements may be merged together into one genus or family, for example, *Pinus* species
331 that are restricted to tropical or subtropical areas in China or ones that grow in boreal forests (Cao et al.,
332 2013; Tian et al., 2017).

333 Along with the pollen assemblages, data repositories also provide chronological information for fossil
334 records. The quality of such chronologies varies strongly with respect to dating methods, calibration and
335 numerical algorithms for determining an age-depth relationship (Blois et al., 2011; Trachsel and Telford,
336 2017). Having accurate and precise chronologies is thus of pivotal importance for reconstructing past
337 climate in order to identify temporo-spatial patterns and therefore in helping to evaluate climate model
338 outputs. The advantage of the fossil pollen dataset used for the reconstruction presented here (i.e.,
339 LegacyPollen 1.0; Herzschuh et al., submitted) is that it has harmonized chronologies (LegacyAge 1.0)
340 along with information about uncertainties as well as related metadata and scripts that allow a
341 customized re-establishment of the chronologies (Li et al., 2022). This, for example, allows the
342 calculation of the temporal uncertainties when presenting reconstruction uncertainties of a specific time-
343 slice.

344

345 **5.2 Modern pollen and climate data sources and LegacyClimate 1.0 quality**

346 Palaeoclimate reconstruction methods such as MAT and WA-PLS rely on extensive collections of
347 modern training data. Designing a robust calibration dataset from modern pollen assemblages is a
348 crucial part of the reconstruction process. A suitable calibration dataset should cover a wide range of
349 climatic and environmental gradients in order to represent an empirical relationship between pollen
350 assemblages and climate (Birks et al., 2010; Chevalier et al., 2020). Like with fossil pollen records, data
351 syntheses and repositories also exist for modern surface pollen data. Most of the records in our modern
352 dataset were compiled from well-established pollen assemblages from North America (Whitmore et al.,
353 2005), Eurasia (Davis et al., 2013 and 2020) and China (Cao et al., 2013; Herzschuh et al., 2019). For
354 fossil pollen records in areas with an insufficient coverage of modern surface pollen samples (e.g.,
355 Central Asia or Western Siberia), it might be difficult to create a calibration dataset that maps the
356 required variety of environmental and climatic gradients and therefore find enough modern analogues



357 for reconstructions with a classification approach such as MAT. Our routine uses the modern pollen data
358 from within a radius of 2000 km around the site of the fossil record. The information provided in the
359 reconstruction metadata including number of modern pollen samples and ranges of reconstructed
360 variables, allows an assessment of the modern dataset used for reconstruction.

361

362 **5.3 Reconstruction method and LegacyClimate 1.0 quality**

363 Climate reconstruction methods all have different strengths and weaknesses. MAT and WA-PLS used
364 in this study heavily rely on extensive collections of modern assemblage data covering diverse climatic
365 and environmental gradients and are applicable on a broad spatial scale. However, both methods may
366 struggle with complex species responses, are sensitive to spatial autocorrelation, can only deal with a
367 certain extent of non-analogous situations and may produce poor results in so-called “quantification
368 deserts” (Chevalier, 2019), where fossil pollen is hardly preserved or nearby modern surface pollen
369 samples are missing (Chevalier et al., 2020). Nonetheless, for reconstructions on a local or regional
370 scale, MAT and WA-PLS are most commonly used in climate reconstructions. The format of the modern
371 and fossil datasets as well as the provided scripts could also be easily adapted to apply to other
372 reconstruction methods such as CREST, a Bayesian approach that combines presence-only occurrence
373 data and modern climatologies to estimate the conditional response of a given taxon to a climate variable
374 (Chevalier et al., 2014 and 2020).

375 Through numerous physical processes that vary with both location and time, temperature and
376 precipitation are interconnected, especially within the extratropical regions (Adler et al., 2008; Trenberth,
377 2011) and thus temperature and precipitation may not be treated as independent variables. Due to the
378 numerical mechanisms in the transfer function, the correlation between both climate variables may
379 reduce the reliability of the reconstructions. This is especially true for regions with a temperature-
380 moisture driven circulation system such as the East Asian Summer Monsoon (EASM) that can heavily
381 affect precipitation patterns in certain regions (Herzschuh et al., 2019). With our tailoring approach we
382 are able to reduce the influence of co-variation of these two climate variables for the reconstruction and
383 increase the number of records that pass a significance level of $p < 0.1$ (Telford and Birks, 2011).



384 **5.4 Potential use of LegacyClimate 1.0**

385 Our fossil pollen synthesis contains records from all over the Northern Hemisphere extratropics and thus
386 can be used to infer spatio-temporal patterns in climate reconstructions that are not only limited to a
387 local or regional scale. Although several hemispheric or global reconstruction studies exist, they have
388 been largely restricted to temperature or have included relatively few records (Marcott et al., 2013;
389 Marsicek et al., 2018; Routson et al., 2019; Kaufman et al., 2020a and 2020b). Our dataset is therefore
390 a valuable addition. It may be used in a multi-proxy approach, synthesizing marine and terrestrial records
391 in order to assess temperature development during the Holocene and can help to highlight possible
392 interdependencies between oceans and land masses and such contribute to the “Holocene conundrum”
393 debate (Liu et al., 2014). Temperature reconstructions from proxy data indicate peak temperatures
394 during the Holocene Thermal Maximum around 6000 years BP followed by a pronounced cooling trend
395 toward the late Holocene (Liu et al., 2014; Bova et al., 2021), which is also visible in our pollen-based
396 reconstructions (Fig. 6). In contrast, climate models simulate a monotonic warming throughout the
397 Holocene, which resulted in the “Holocene conundrum” debate (Liu et al., 2014). Temperature
398 reconstructions are often derived from sea-surface temperatures as either mean annual temperatures
399 (Birks, 2019; Bova et al., 2021) or global mean surface temperature (Marcott et al., 2013; Marsicek et
400 al., 2018; Kaufman et al., 2020a and 2020b). However, it is argued that proxy-based climate
401 reconstructions are seasonally biased and therefore might be the reason for the observed proxy-model
402 divergence (Liu et al., 2014; Rehfeld et al., 2016; Bova et al., 2021). In this respect, it might help that
403 we provide T_{july} along with T_{ann} reconstructions, which provides the opportunity to assess seasonal
404 impacts on the reconstruction.

405 So far, reconstructions of precipitation have not been implemented on a hemispheric scale. The
406 interconnection between temperature and precipitation (Trenberth, 2011) and its spatio-temporal
407 variation across the Northern Hemisphere is therefore an important aspect of evaluating climate models
408 (Wu et al., 2013; Hao et al., 2019; Herzs Schuh et al., submitted). A broad-scale quantitative reconstruction
409 of temperature and precipitation would therefore be of great value for evaluating transient model runs
410 performed by climate models such as TraCE 21k (He, 2010).

411 **6 Data and code availability**

412 The compilation of reconstructed T_{july} , T_{ann} , and P_{ann} , is open access and available at PANGAEA
413 (<https://doi.pangaea.de/10.1594/PANGAEA.930512>; in the “*Other version*” section; Herzs Schuh et al.,



414 2021). The dataset files are stored in machine-readable data format (.CSV), which are already separated
415 into Western North America, Eastern North America, Europe, and Asia for easy access and use.

416 The R code to run the reconstructions for single sites is available at Zenodo
417 (<https://doi.org/10.5281/zenodo.5910989>; Herzs Schuh et al., 2022) including harmonized open-access
418 modern and fossil pollen datasets so that customized reconstructions can be easily established.

419

420 **Author contributions.** UH designed the study design and reconstruction dataset. CL and TB compiled
421 the metadata and the harmonized pollen dataset. TB wrote the R scripts and ran the analyses under the
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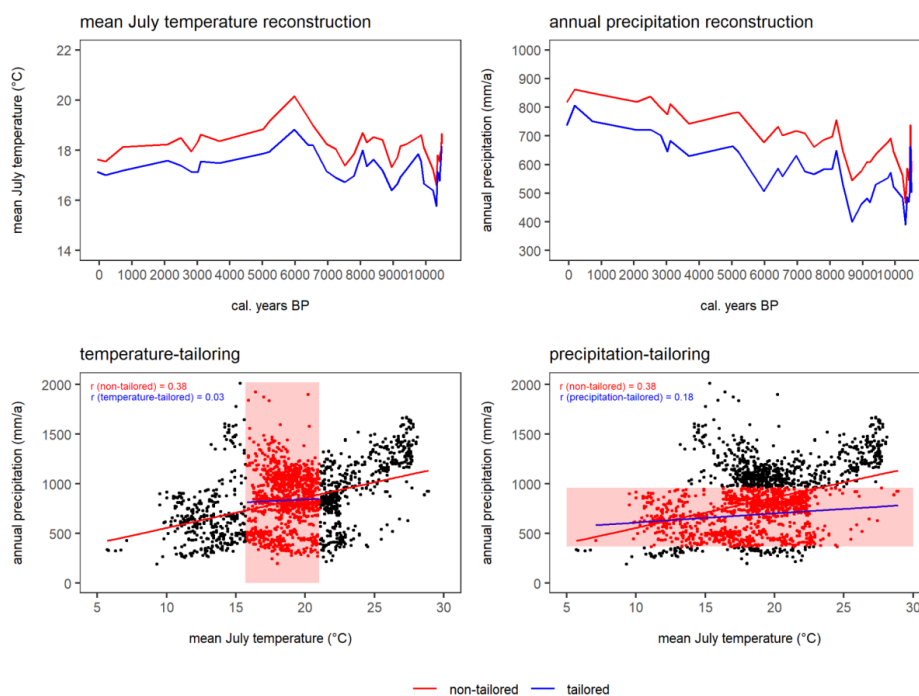
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640

641 Appendix Figures



642



643 **Appendix Figure 1.** Example to illustrate the effect of tailoring the modern dataset for the location
644 “Yellow Dog Pond” in Eastern North America. Upper part: reconstruction of T_{July} and P_{ann} with WA-PLS
645 (red) and WA-PLS_tailored (blue); lower part: correlation of T_{July} and P_{ann} in the modern dataset and the
646 effect of tailoring the modern dataset (indicated with the red box). Correlations are given for non-tailored
647 (red) and tailored (blue) data.