



1 **Ice core chemistry database: an Antarctic compilation of**
2 **sodium and sulphate records spanning the past 2000 years.**

3 Elizabeth R. Thomas¹, Diana O. Vladimirova¹, Dieter Tetzner¹, B. Daniel Emanuelsson¹,
4 Nathan Chellman², Daniel A. Dixon³, Hugues Goosse⁴, Mackenzie M. Grieman⁵, Amy C.F.
5 King¹, Michael Sigl⁶, Danielle G Udy⁷, Tessa R. Vance⁸, Dominic A. Winski³, V. Holly L.
6 Winton⁹, Nancy A.N. Bertler^{9,10}, Akira Hori¹¹, Chavarukonam.M Laluraj¹², Joseph R.
7 McConnell², Yuko Motizuki¹³, Kazuya Takahashi¹³, Hideaki Motoyama¹⁴, Yoichi Nakai¹³,
8 Franciele Schwannck¹⁵, Jefferson Cardia Simões¹⁵, Filipe G. L. Lindau¹⁵, Mirko Severi¹⁶, Rita
9 Traversi¹⁶, Sarah Wauthy¹⁷, Cunde Xiao¹⁸, Jiao Yang¹⁹, Ellen Mosely-Thompson²⁰, Tamara
10 V. Khodzher²¹, Ludmila P. Golobokova²¹, Alexey A. Ekaykin²²

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12 ¹Ice Dynamics and Paleoclimate, British Antarctic Survey, High Cross, Madingley Road, Cambridge,
13 CB3 0ET, UK

14 ²Division of Hydrologic Sciences, Desert Research Institute, Reno, NV, 89512, USA

15 ³Climate Change Institute, University of Maine, 5790 Bryand Global Science Center, Orono, ME,
16 04469, USA.

17 ⁴Earth and Life Institute, Universite catholique de Louvain, Place Pasteur 3, 1348 Louvain-la-Neuve,
18 Belgium

19 ⁵Department of Chemistry, Reed College, 3203 Woodstock Blvd., Portland, Oregon, 97202, USA

20 ⁶Climate and Environmental Physics (CEP), Physics Institute & Oeschger Centre for Climate Change
21 Research (OCCR), University of Bern, Sidlerstrasse 5, 3012 Bern, Switzerland

22 ⁷Institute for Marine and Antarctic Studies, University of Tasmania, 20 Castray Esplanade, Battery
23 Point TAS 7004, Australia

24 ⁸Australian Antarctic Program Partnership, Institute for Marine & Antarctic Studies, University of
25 Tasmania, Hobart, Australia

26 ⁹Antarctic Research Centre, Victoria University of Wellington, Kelburn Parade, Kelburn, Wellington
27 6021, New Zealand

28 ¹⁰National Ice Core Facility, GNS Science, 30 Gracefield Rd, Gracefield 5040, New Zealand

29 ¹¹Kitami Institute of Technology, 090-8507, Japan

30 ¹²National Centre for Polar and Ocean Research (NCPOR), Ministry of Earth Sciences, Vasco-da
31 Gama, Goa 403804, India

32 ¹³RIKEN Nishina Center for Accelerator-Based Science, 2-1 Hirosawa, Wako, Saitama 351-0198,
33 Japan

34 ¹⁴National Institute of Polar Research, Tachikawa, Tokyo 190-8518, Japan

35 ¹⁵Centro Polar e Climático, Universidade Federal do Rio Grande do Sul, Porto Alegre, 91501-970,
36 Brazil

37 ¹⁶Chemistry Dept. "Ugo Schiff", University of Florence, 50019, Sesto F.no, Florence, Italy.



38 ¹⁷Laboratoire de Glaciologie, Department Geosciences, Environnement et Societe, Universite Libre de
39 Bruxelles, 1050 Brussels, Belgium

40 ¹⁸State Key Laboratory of Earth Surface Processes and Resource Ecology, Beijing Normal University,
41 China

42 ¹⁹State Key Laboratory of Cryospheric Science, Northwest Institute of Eco-Environment and
43 Resources, Chinese Academy of Sciences, Lanzhou 730000, China

44 ²⁰Byrd Polar and Climate Research Center, The Ohio State University, 1090 Carmack Rd. Columbus
45 OH 43210 USA

46 ²¹ Limnological Institute of Siberian Branch of the Russian Academy of Sciences), Irkutsk, 664033,
47 Russia

48 ²² Arctic and Antarctic Research Institute), 38 Bering st, St Petersburg, 199397, Russia

49 *Correspondence to:* Elizabeth R. Thomas (lith@bas.ac.uk)

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74 **Abstract.** Changes in sea ice conditions and atmospheric circulation over the Southern Ocean play an important
75 role in modulating Antarctic climate. However, observations of both sea ice and wind conditions are limited in
76 Antarctica and the Southern Ocean, both temporally and spatially, prior to the satellite era (1970 onwards). Ice
77 core chemistry data can be used to reconstruct changes over annual, decadal, and millennial timescales. To
78 facilitate sea ice and wind reconstructions, the CLIVASH2k working group has compiled a database of two
79 species, sodium [Na^+] and sulphate [SO_4^{2-}], commonly measured ionic species. The database contains records
80 from 105 Antarctic ice cores, containing records with a maximum age duration of 2000 years. An initial filter
81 has been applied, based on evaluation against climate observations, to identify sites suitable for reconstructing
82 past sea ice conditions, wind strength, or atmospheric circulation.

83

84

85 1 Introduction

86 Changes in wind strength and atmospheric circulation, notably the positive phase of the Southern Annular Mode
87 (SAM), have been related to increased Antarctic snowfall (Thomas et al., 2017; Thomas et al., 2008; Medley
88 and Thomas, 2019) and the widespread warming in the Antarctic Peninsula (Turner et al., 2016; Thomas et al.,
89 2009) and West Antarctica during the 20th century. Contemporaneously, Antarctic sea ice is also undergoing
90 significant change. Despite model predictions of a homogeneous decline (Roach et al., 2020), total Antarctic sea
91 ice cover has increased since 1970 (Zwally et al., 2002; Turner et al., 2009). With more recent periods of abrupt
92 decline in 2016, (Meehl et al., 2016) and 2022 (Turner et al., 2022).

93 Our understanding of winds, atmospheric circulation and sea ice is hampered by both the lack of observations
94 prior to the instrumental period (~1970s onwards) and uneven spatial coverage of paleoclimate records (Jones et
95 al., 2016; Thomas et al., 2019). Data-model intercomparison and data synthesis studies have demonstrated the
96 value of large datasets in reconstructing climate and sea ice over decadal to centennial scales (Dalaiden et al.,
97 2021; Fogt et al., 2022). To meet the need for Antarctic-wide, spatially dense, and intercomparable atmospheric
98 circulation and sea ice records, we propose the use of chemical species routinely measured in ice cores.

99 Sodium [Na^+], from sea salt aerosol, has been proposed as a proxy for past sea ice extent (SIE)
100 (Waisdivideprojectmembers et al., 2013; Severi et al., 2017; Wolff et al., 2006; Winski et al., 2021). The sea
101 salt component of [Na^+] arises from both sea ice and open water and the relationship between [Na^+] and sea ice
102 varies between sites (Sneed et al., 2011). High winds mobilize [Na^+] from the sea ice surface, either in frost
103 flowers or brine-soaked snow (Huang and Jaeglé, 2017; Frey et al., 2020). The [Na^+] reaching the ice core sites
104 is dependent on both the distances from the source, either sea ice or open ocean, and the meteorological
105 conditions (Minikin et al., 1994). [Na^+] is therefore a valuable tracer for marine-air mass advection and changes
106 in atmospheric circulation (Dixon et al., 2004; Mayewski et al., 2017).

107 Sulphate [SO_4^{2-}] is formed in the atmosphere as secondary aerosol following volcanic and anthropogenic
108 sulphur dioxide [SO_2] gas emissions. [SO_4^{2-}], together with methane sulphonic acid [MSA], is the main
109 atmospheric sulphur compound formed from ocean-derived dimethylsulfide (DMS) (Gondwe et al., 2003). In
110 the southern hemisphere, marine biogenic emissions dominate the total sulphur budget (Delmas et al., 1982;
111 Legrand and Mayewski, 1997; McCoy et al., 2015). Sulphur can significantly impact cloud albedo and new
112 particle formation (Brean et al., 2021). The sea salt fraction of [SO_4^{2-}] is largest at coastal and low elevation sites
113 (Dixon et al., 2004). The non-sea salt fraction, also referred to as excess [SO_4^{2-}] (hereafter referred to as xs
114 [SO_4^{2-}]), can be estimated based on the relationship with [Na^+] (e.g., $\text{xs} [\text{SO}_4^{2-}] = [\text{SO}_4^{2-}] - 0.25[\text{Na}^+]$) (O'Brien et
115 al., 1995). Excess [SO_4^{2-}] has been shown to correlate with SIE at some ice core sites (Dixon et al., 2004; Sneed
116 et al., 2011). The background xs [SO_4^{2-}] source, from marine biogenic deposition, is superimposed by sporadic
117 volcanic deposition of [SO_4^{2-}] providing an excellent reference horizon for dating Antarctic ice cores (Dixon et
118 al., 2004; Sigl et al., 2014; Plummer et al., 2012). At low elevation and coastal sites, where background biogenic
119 sources are high, it is not always possible to identify volcanic peaks (Emanuelsson et al., 2022; Tetzner et al.,
120 2021b). In this study, [SO_4^{2-}] provides a dual function: 1) as a potential proxy for SIE and 2) as a stratigraphic
121 age marker to validate submitted age-scales and subsequently align ice-core chronologies onto a common
122 chronology.

123 1.1. The CLIVASH2k chemistry database.



124 CLIVASH2k (CLimate Variability in Antarctica and the Southern Hemisphere over the past 2000 years) is a
125 project of the Past Global Changes (PAGES) 2k network. The CLIVASH2k database is the latest in a series of
126 community-led paleoclimate data synthesis efforts endorsed by PAGES (Kaufman et al., 2020; McGregor et al.,
127 2015; McKay and Kaufman, 2014; Tierney et al., 2015; Thomas et al., 2017; Stenni et al., 2017; Konecky et al.,
128 2020). The aim of this study is to focus on two primary species, sodium, and sulphate, as they are routinely
129 measured in ice cores and have potential links with either sea ice or atmospheric circulation. The time window
130 of the last 2000 years has been selected to cover both natural and anthropogenic changes.

131 Two main features distinguish the CLIVASH2k data compilation from previous PAGES synthesis: 1) the data
132 included are not limited to previously published records, and 2) the data comprise two distinct chemical species
133 which do not have a well-established relationship with climate (beyond the episodic sources of $[\text{SO}_4^{2-}]$ noted
134 above).

135 Calls for participation in CLIVASH2k activities were widely distributed, ensuring a representative cross section
136 of scientists from various disciplines, geographic regions, and career stage. The targeted species to target and the
137 selection criteria were decided at several open discussion stages, followed by updates to the CLIVASH2k
138 mailing list and distributed via PAGES monthly updates.

139 2. Methods

140

141 2.1. Resolution and duration.

142 The target time-period for the database is the last 2000 years. Records of any duration could be submitted within
143 this time-period. These records could be from snow-pits and firn cores, spanning just a few seasons to years.
144 Data were requested at the highest resolution available and converted to annual averages (January – December).
145 Years with missing data were included, providing a threshold of half a year of data was achieved.

146

147 2.2. Age-scales.

148 Most records within this time-period have been annually dated, based on the seasonal deposition of distinct
149 chemical species (including sodium, sulphate, and sulphur). The longer records, those spanning the last 500-
150 2000 years, have been synchronized previously (Sigl et al., 2014) or within this project on the WD2014 age-
151 scale (Sigl et al., 2016) or have age-scales that are broadly consistent with WD2014 (Plummer et al., 2012). This
152 new chronology is constrained by the 774 CE cosmogenic (i.e. ^{10}Be) anomaly, and is consistent with
153 dendrochronology (Büntgen et al., 2018) and ice core chronologies from Greenland (Sigl et al., 2015). The
154 WD2014 age-scale is recommended because it is consistent with the forcings applied in PMIP4/CMIP6 model
155 simulations (Jungclauss et al., 2017). Age transfer functions can now be linked to other PAGES2k
156 reconstructions and individual ice cores. There are two exceptions, Plateau Remote and DT401 (both very low
157 accumulation sites in the interior of east Antarctica), which differ from WD2014 prior to 1000 AD and cannot
158 be confidently synchronized. The third exception is partly unpublished data from the Vostok vicinity, which
159 were dated using the snow accumulation rate and volcanic age markers (this study and (Ekaykin et al., 2014).

160

161 2.3. Peer review and publications.

162 Unlike previous PAGES 2k compilations, the CLIVASH2k database was not constrained by the need for
163 records to be published and peer reviewed. This decision arose based on the limited number of published
164 chemistry records available and the desire to maximise the records. Published records were submitted along
165 with their original citation; unpublished records were listed as “This study”, with the data contributor included
166 as a co-author.

167

168 2.4. Analytical methods.

169 Both the ionic and elemental forms of sodium ($[\text{Na}]$ and $[\text{Na}^+]$) and sulphur ($[\text{S}]$ and $[\text{SO}_4^{2-}]$), respectively, were
170 accepted as part of the CLIVASH2k data call. Several analytical techniques are used to measure $[\text{Na}^+]$, $[\text{S}]$ and
171 $[\text{SO}_4^{2-}]$ in ice cores. Ionic $[\text{Na}^+]$ and $[\text{SO}_4^{2-}]$ are typically measured by ion chromatography (IC), while elemental
172 Na and S are generally measured by inductively coupled plasma mass spectrometry (ICP-MS). Unlike IC, which
173 measures the soluble fraction, ICP-MS techniques measure the total elemental concentration of both the
174 dissolved and particulate fraction of the element. While continuous ICP-MS measurements of certain species



175 may require correction for under-recovery, Na and S are typically fully recovered during continuous
176 measurements (Arienzo et al., 2019). Previous comparisons of analytical methods show excellent agreement of
177 [Na] in ice cores measured using IC and ICP-MS methods e.g. (Grieman et al., 2022). This agreement suggests
178 that the ionic and elemental forms reported in the database can be directly compared.

179
180 Biogenic atmospheric emissions of organic [S] species, mainly dimethyl sulfide (DMS), are a major contributor
181 to the [S] in the Antarctic snow. In the marine atmosphere DMS is oxidized to [MSA⁻] and [SO₄²⁻], which are
182 eventually deposited on the polar ice sheets. The ICP-MS technique measures total [S] in ice cores, which
183 includes [S] contained [MSA⁻]. In contrast, the IC technique solely quantifies [S]. If total [S] and [MSA⁻] are
184 both analysed on the same ice core, the [MSA⁻] contribution can be subtracted (Cole-Dai et al., 2021). However,
185 continuous [MSA⁻] measurements are scarce over Antarctica (Thomas et al., 2019) and the long-term variability
186 of both [MSA⁻] and [SO₄²⁻] is very small during the common era (Legrand et al., 1992; Saltzman et al., 2006).
187 Thus, we applied a consistent transformation across all sites. We multiplied elemental [S] (32 g mol⁻¹) from
188 ICP-MS measurements with three to convert to the equivalent [SO₄²⁻] (96 g mol⁻¹) without applying corrections
189 for MSA contributions. To aid ease of comparison, all [S] has been converted to [SO₄²⁻], in the database and will
190 be referred to only as [SO₄²⁻] in the data description.

191 **2.5. Flux vs concentration.**

192 [Na⁺] and [SO₄²⁻] in ice cores are generally reported as a concentration. Concentration can be converted to a
193 deposition flux, provided that the snow accumulation rate is known. The CLIVASH2k database includes both
194 concentrations and fluxes, when available. Flux estimates from ice cores combine both wet and dry deposition,
195 of which the contribution of these two depositional modes varies across Antarctica with elevation and distance
196 from the source (Wolff, 2012).

197

198 **2.6. Establishing the sea salt and non-sea salt component.**

199 There are various methods of calculating the sea salt (ss) and excess (xs) components of an ice core chemistry
200 record. The most-common method, as mentioned above, is to assume 100% of the [Na⁺] comes from the ocean.
201 Then [Na⁺] can be treated as a marine reference species and the ss fraction of all other chemical species can be
202 calculated based upon a mean ocean water elemental abundance reference value (e.g. (Lide, 2005). If [Na⁺] is
203 suspected of not being of marine origin, alternative methods of calculating the ss chemical fraction may be
204 employed. For example, one may apply a standard sea-water ratio of 30.61 [Na⁺], 1.1 [K⁺], 3.69 [Mg²⁺], 1.16
205 [Ca²⁺], 55.04 [Cl⁻] and 7.68 [SO₄²⁻] to the ion concentrations in each sample (Holland, 1978). Several studies
206 have shown that frost flowers are depleted in [SO₄²⁻] relative to [Na⁺]. This produces a ssSO₄²⁻ value which is
207 slightly higher than it should be for sites near the coast (Rankin et al., 2002; Rankin et al., 2000). Unfortunately,
208 not all studies accurately measure a wide suite of chemical species. Therefore, in this study we have assumed
209 [Na⁺] to be the primary marine species and calculated xs [SO₄²⁻] according to the following ratio: [xsSO₄²⁻] =
210 [SO₄²⁻] - 0.25[Na⁺] (O'Brien et al., 1995). Other ratios may be more suitable for coastal sites, but for consistency
211 we have applied the same ratio to all records reported in the database.

212

213 **2.7. Data validation and recommendations**

214 The two main uncertainties in the data presented arise from 1) chronological controls and 2) analytical errors.
215 As discussed in section 2.2, all records have been synchronised to a common age-scale (WD2014). Thus, when
216 using the entire database, we recommend using an error estimate of ±2 years, for records younger than 500
217 years, increasing to a conservative error estimate of ±5 years for records extending to 2000 years. This is the
218 maximum uncertainty estimate for the WD2014 age-scale at 2,500 years (Sigl et al., 2015). However, we note
219 that for individual records in this database the published error estimates are as low as ±1 year (e.g., Emanuelsson
220 et al., 2022). When using individual records we recommend using the published error estimate for that record.

221 Analytical precision varies between instruments and laboratories. We recommend applying a 1 standard error
222 (σ) to the data to account for analytical errors.

223 The [Na⁺] and [SO₄²⁻] data is an accurate representation of either concentration or flux at a certain site.
224 However, how this relates to regional deposition is not well constrained. While we can account for the
225 uncertainty in analytical precision and dating error, we cannot define the signal to noise ratio associated with



226 small scale post-depositional process. For example, wind redistribution or the impact of local orography. The
 227 regional climate and signal to local noise has been investigated for stable water isotopes in Antarctica (e.g.,
 228 Munch et al., 2018), however, a detailed investigation of $[\text{Na}^+]$ and $[\text{SO}_4^{2-}]$ is lacking. One of the main
 229 limitations, which this database will address, has been the lack of available data. We thus encourage database
 230 users to investigate the regional signal by averaging records to reduce the signal to noise ratio. In this case, we
 231 recommend using the standard error propagation procedure for averaging for example the square root of the sum
 232 of variances of individual records divided by the number of the records.

233 Ice cores provide the only record of $[\text{Na}^+]$ and $[\text{SO}_4^{2-}]$ deposition in Antarctica, and therefore, validation against
 234 reference datasets is also not possible. While progress has been made using chemical transport models to
 235 represent the deposition of sea salts in Greenland (Rhodes et al., 2018), the period examined is very short
 236 (annual to decadal) and has currently not been applied to Antarctica. This database will provide much needed
 237 data for any future model validation. However, currently it means there are no independent data products to
 238 validate our $[\text{Na}^+]$ and $[\text{SO}_4^{2-}]$ records against.

239

240 3. Data records

241 A total of 117 records were submitted, representing 105 individual ice core sites (Fig. 1). In some locations,
 242 duplicate analysis or updated versions were submitted (e.g. EPICA Dome C). All submitted records have been
 243 included in the database. The number of records submitted is summarised in Table 1. The full list of records,
 244 their location, elevation, duration, and reference are presented in appendix A (table A1).

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247

Table 1. Summary of records submitted to the CLIVASH2k database.

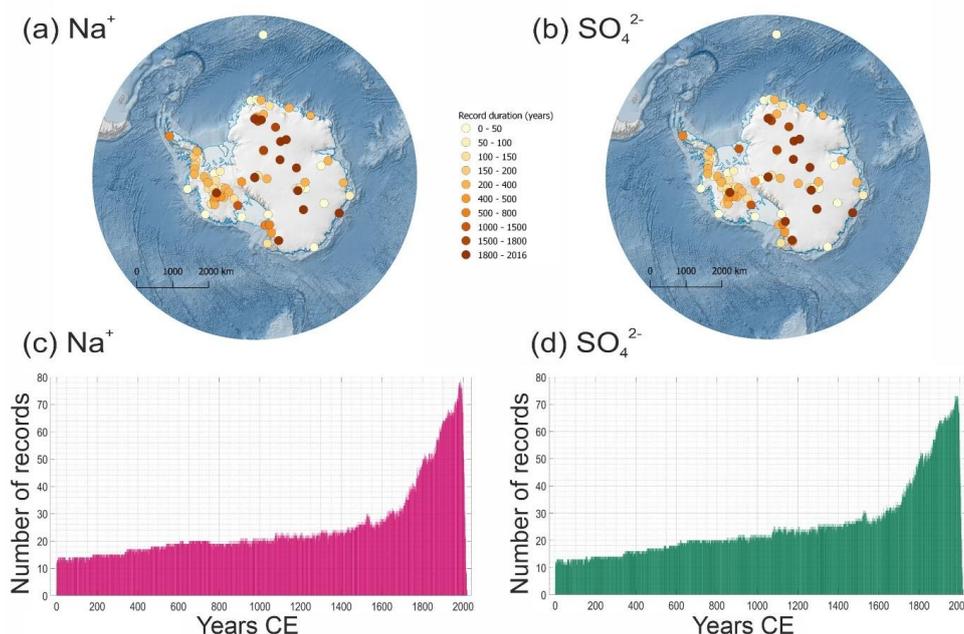
	$[\text{Na}^+]$		$[\text{SO}_4^{2-}]$		xs $[\text{SO}_4^{2-}]$	
	Concentration	Flux	Concentration	Flux	Concentration	Flux
Total records	105	68	103	67	95	61
Duplicates	15	4	13	2	11	2
Total ice core sites	90	64	90	65	84	59

248

249 3.1. Geographical and temporal coverage.

250 There is reasonable spatial coverage across Antarctica, with the largest density of records in West Antarctica
 251 (Figs. 1a & 1b). In East Antarctica, notable data voids include Coats Land, Enderby and Kemp Land, Wilkes
 252 Land and Terra Adelie. There is a notable absence of long records from the Antarctic Peninsula. Despite the
 253 high density of records in West Antarctica, high snow accumulation in this region results in most of these
 254 records only spanning the last few decades or centuries.

255 The longer duration records (>1000 years) are predominantly found in the central East Antarctic plateau, while
 256 most higher snow accumulation coastal sites cover shorter timescales (Figs. 1a & 1b). The most recent year in
 257 the record peaks in the late 1990s, when the highest number of cores were drilled (Figs. 1c & 1d). Only eleven
 258 records span the full 2000 years.



259

260 **Figure 1.** Spatial and temporal coverage of records in the CLIVASH2k database. Map of ice core locations with
261 (a) $[\text{Na}^+]$, and (b) $[\text{SO}_4^{2-}]$ records. Colour coded based on record duration (number of years). The number of (c)
262 $[\text{Na}^+]$ and (d) $[\text{SO}_4^{2-}]$ records as a function of the years (CE) covered.

263

264 3.1.1. Technical validation

265 To facilitate the scientific usability of this database, we have evaluated each record in terms of its relationship
266 with key climate parameters during the observational period (1979- 2019). Given their varying temporal ranges
267 (Fig. 1), not all the records span the full satellite period. Thus, correlations are based on the largest number of
268 years available within this period. Although the database includes short records, for the data interpretation step,
269 we have only included records that have at least ten years of overlap with the satellite and reanalysis climate
270 data. Duplicate records (including updated versions and different analytical approaches) are included in the data
271 interpretation step and interpreted as individual records.

272

273 The objective of this climatological comparison is to provide a first level filter for the database. Thus enabling
274 database users to quickly search for sites that exhibit a direct and dynamically logical relationship with sea ice
concentration (SIC), wind conditions and atmospheric circulation to facilitate future investigations.

275

276 All of the records were also correlated using ERA5 meteorological parameters (Hersbach et al., 2020), the fifth
277 generation European Centre for Medium Range Weather Forecast (ECMWF) atmospheric reanalysis data. These
278 parameters include 500-hPa geopotential height (Z500), meridional winds (v) and zonal winds (u) both at the
279 850-hPa level. The 850 hPa level was chosen to represent surface winds (relevant for sea ice reconstructions),
280 while the 500 hPa was chosen to capture larger-scale circulation across both high and low elevation sites. All
281 correlations were performed on de-trended annual average data (January – December) to correspond with the
282 annually-resolved ice core records and corrected for autocorrelation. All of the records were correlated with SIC
283 from the National Snow and Ice Data Centre (NSIDC) Nimbus-7 SMMR and DMSP SSM/I-SSMIS Passive
Microwave Data version 1 (Cavalieri et al., 1997).

284

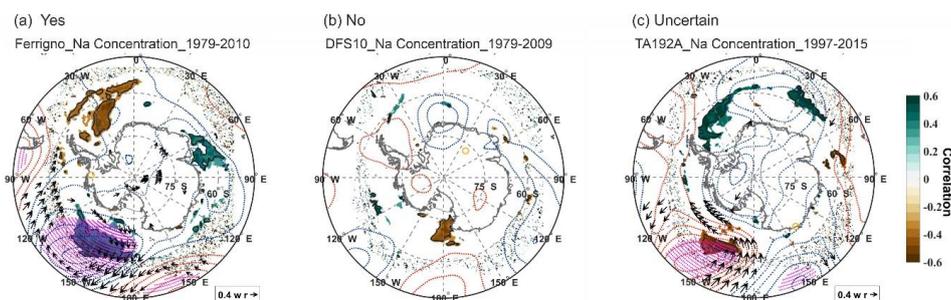
285



286 **4. Data interpretation**
287 **4.1. Identifying sites that correlate with sea ice and atmospheric circulation**

288 An example of the data interpretation output is presented in Figure 2. For consistency, correlations were
289 performed with climate variables across all longitudes in the southern hemisphere south of 50°S. This approach
290 has the potential to generate spurious results or correlations in regions that are physically unrelated to the site
291 (e.g., Fig. 2b). Therefore, each record was individually evaluated by an expert (hereafter the “interpretation
292 team”) to establish if the correlations observed can be attributed to a realistic source region and transport
293 mechanism. Sites with a clear connection or absence of connection agreed by more than one interpreter were
294 marked as either “yes” or “no” (Figs. 2a & 2b). Sites where the transport mechanism was less clear, or there was
295 a disagreement between interpreters were listed as “uncertain” (Fig. 2c).

296



297

298 **Figure 2.** Example correlation plots evaluated by the data interpretation team. (a) Yes example, correlation
299 observed between all three parameters. (b) No example, no significant correlation observed with any parameters.
300 In this example, a significant correlation with SIC at a distant location is likely an auto-correlation artefact. (c)
301 Uncertain example, the transport mechanism could not be verified by the interpretation team. Yellow open circle
302 indicates ice core location. Coloured shading indicates positive (green) and negative (brown) correlations with
303 SIC (data from NSIDC), solid black line correlations significant at the 5% level. Correlations with winds
304 (arrows) composed of u850 and v850 (ERA 5). Dashed red and blue contours represent positive (red) and
305 negative (blue) correlations with geopotential height at 500 hPa (ERA5), pink hatching is significant at the 5%
306 level. Plot titles labelled as “Site name_species_years for correlation”.

307

308 In the following sections, we only refer to records that exhibited a correlation that is statistically significant at
309 the 5% level ($p < 0.05$) (hereafter referred to as significant). For sites to be identified as having a relationship
310 with either SIC, atmospheric pressure (z500) or winds (u850 or v850), they had to be supported by a valid
311 transport mechanism or source region as evaluated by the data interpretation team (Fig. 2). We have not applied
312 a uniform cut-off size for the area of correlation or specified a minimum or maximum distance from the source
313 region, as these features will be site specific. For example, a low elevation coastal site (e.g., Sherman Island)
314 may only capture local changes in sea ice that will appear as a small area of correlation on the map (e.g.,
315 (Tetzner et al., 2021a) while a central Antarctic site (e.g., South Pole) might be influenced by long-range air-
316 masses and thus capture changes in sea ice from a relatively distant source region e.g., (Winski et al., 2021).

317 The database contains more concentration records than flux records. Thus, in the data interpretation we
318 presented both the total number of sites, and the proportion of sites, that exhibit a significant correlation with
319 meteorological parameters. The total number of eligible records for each species is shown in Table 3. The
320 spatial distribution of records is presented in figures 3, 4 and 5.

321

322 **Table 3.** Summary of the number of records that display a significant correlation (5% level) with SIC, wind
323 fields (meridional (v850) and zonal (u850)), and geopotential height (z500). The total records available for the
324 data interpretation step is shown for each species. This includes all records with more than 10-years overlap



325 with the instrumental period (1979-2018) and includes duplicates. Brackets indicate the number of sites marked
 326 as “uncertain”. The percentage of records shown in italics underneath to account for the varying sample size.

327

Variable	[Na ⁺]	Na ⁺ Flux	[SO ₄ ²⁻]	SO ₄ ²⁻ Flux	xs [SO ₄ ²⁻]	xs SO ₄ ²⁻ Flux
Total records	88	65	84	61	81	59
SIC	69 (6) <i>78 %</i>	56 (4) <i>86 %</i>	60 (6) <i>71 %</i>	40 (5) <i>66 %</i>	68 (5) <i>84 %</i>	42 (2) <i>71 %</i>
Wind (v850 or u850)	63 (3) <i>72 %</i>	48 (4) <i>74 %</i>	54 (8) <i>64 %</i>	39 (3) <i>64 %</i>	56 (3) <i>69 %</i>	40 (3) <i>68 %</i>
Geopotential Height (z500)	47 (2) <i>53 %</i>	43 (3) <i>66 %</i>	38 (6) <i>45 %</i>	26 (3) <i>43 %</i>	40 (6) <i>49 %</i>	23 (3) <i>39 %</i>

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330

4.2. Sodium (concentration and flux)

331

A total of 69 (out of 88) [Na⁺] sites exhibit a correlation with SIC, with an additional six records marked as
 332 “uncertain” (Table 3). Fifty-Six (out of 65) records are correlated with SIC when using Na⁺ flux, with an
 333 additional four sites marked as uncertain. This reflects the smaller number of flux records submitted to the
 334 database. Proportionally, more records are correlated with SIC when using flux than concentration (78 %
 335 compared to 72 %).

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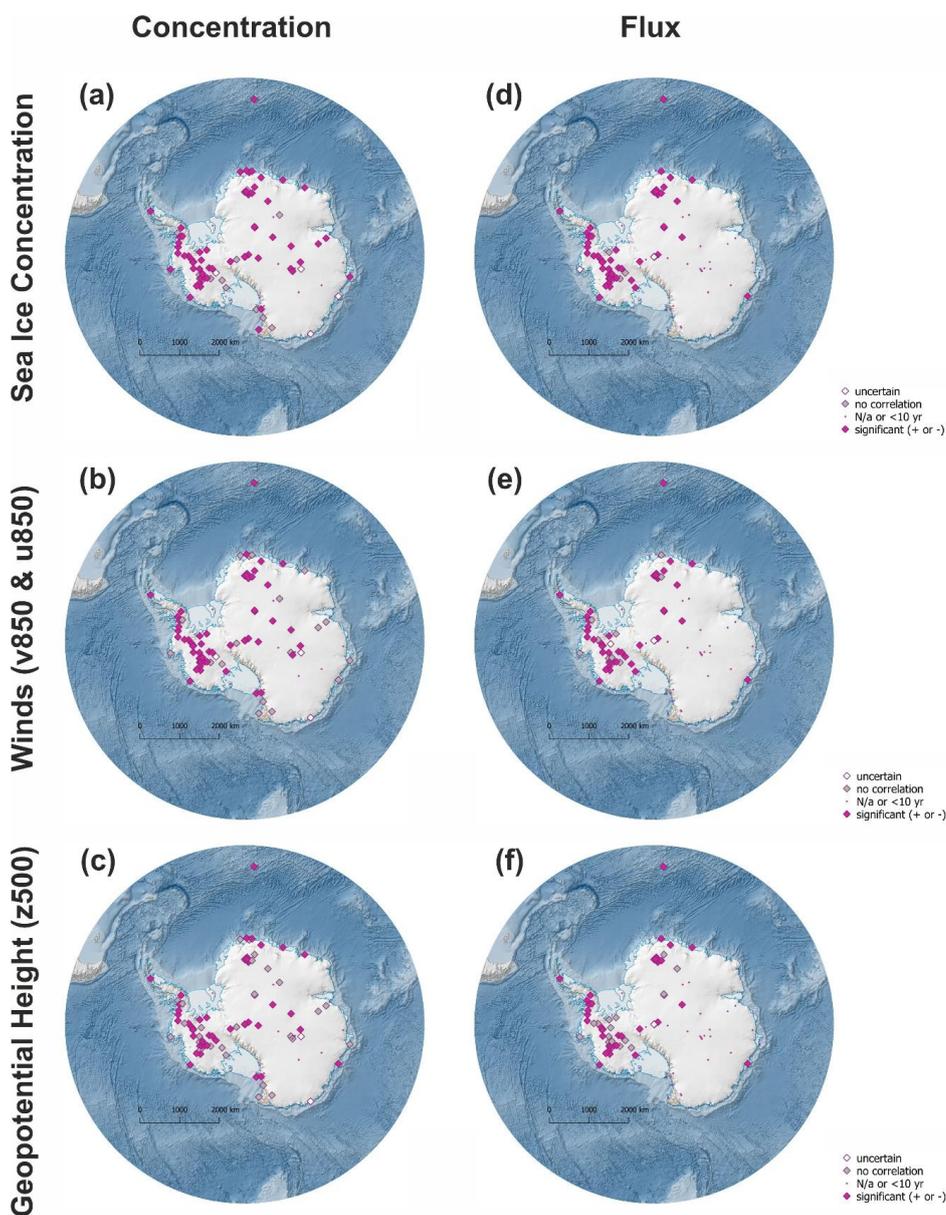
A total of 63 (out of 88) [Na⁺] records exhibit a significant correlation with the wind fields (v850 and u850).
 338 While an additional four records were marked as uncertain. When using Na⁺ flux 48 (out of 65) records
 339 correlated with winds, with four records marked as uncertain. A higher proportion of records (74 % compared
 340 with 72 %) correlated with winds when using flux.

341

342

A total of 47 (out of 88) [Na⁺] sites exhibit a significant correlation with geopotential height. While an
 343 additional two records are marked as uncertain. The number of correlations with geopotential height is 43 (out
 344 of 65) when using Na⁺ flux, with an additional three sites marked as uncertain. A higher proportion of records
 345 (66 % compared with 53 %) correlated with atmospheric circulation when using flux.

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Figure 3 – Geographical distribution of $[Na^+]$ records (left column) which exhibit a statistically significant ($p > 0.05$) correlation with (a) SIC, (b) winds (v850 and u850) and (c) geopotential height (z500). Compared with the geographical distribution of Na flux record (right column) which exhibit a statistically significant ($p > 0.05$) correlation with (d) SIC, (e) winds (v850 and u850) and (f) geopotential height (z500). Pink diamonds are locations with a significant correlation either positive or negative; grey diamonds are sites with no correlation, open diamonds are uncertain. Dots indicate ice core locations that are in the database but either are less than 10 years in length (or overlap with the instrumental period) or sites which failed to generate any correlations with parameters tested.

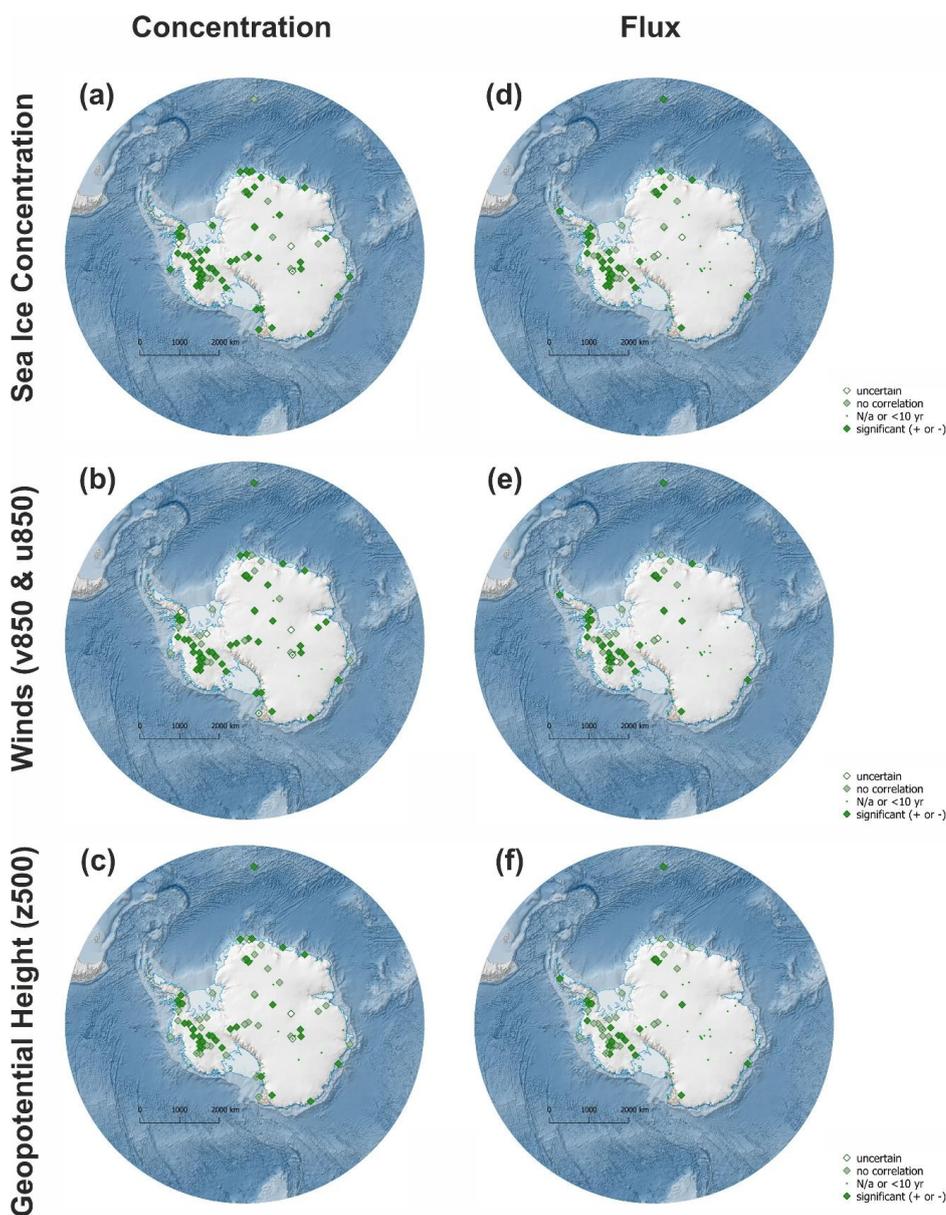


359 **4.3. Sulphate (concentration and flux)**

360 A total of 60 (out of 84) $[\text{SO}_4^{2-}]$ records display a correlation with SIC, with six additional records marked as
361 uncertain (Table 3). When using SO_4^{2-} flux, 39 (out of 61) records correlated with SIC, with an additional five
362 records marked as uncertain. A slightly higher proportion of records (71 % compared with 64 %) correlated with
363 SIC when using flux.

364 Fifty-four $[\text{SO}_4^{2-}]$ records (out of 84) are correlated with winds (v850 and u850), with eight additional records
365 marked as uncertain. This is compared to 39 records (out of 61), and three additional records marked as
366 uncertain, that are correlated with winds when using SO_4^{2-} flux. The proportion of records correlated with winds
367 (64 %) is the same when using either flux or concentration.

368 A total of 38 (out of 84) $[\text{SO}_4^{2-}]$ records are correlated with geopotential height, with six additional records
369 marked as uncertain. This is compared with 26 records (out of 59) when using flux, with three marked as
370 uncertain. A slightly higher proportion of records (45 % compared with 43 %) are correlated with atmospheric
371 circulation when using flux.



372

373 **Figure 4** – Geographical distribution of $[\text{SO}_4^{2-}]$ records (left column) which exhibit a statistically significant
374 ($p > 0.05$) correlation with (a) SIC, (b) winds (v850 and u850) and (c) geopotential height (z500). Compared with
375 the geographical distribution of SO_4^{2-} flux record (right column) which exhibit a statistically significant ($p > 0.05$)
376 correlation with (d) SIC, (e) winds (v850 and u850) and (f) geopotential height (z500). Green diamonds are
377 locations with a significant correlation either positive or negative; grey diamonds are sites with no correlation,
378 open diamonds are uncertain. Dots indicate ice core locations that are in the database but either are less than 10
379 years in length (or overlap with the instrumental period) or sites which failed to generate any correlations with
380 parameters tested.

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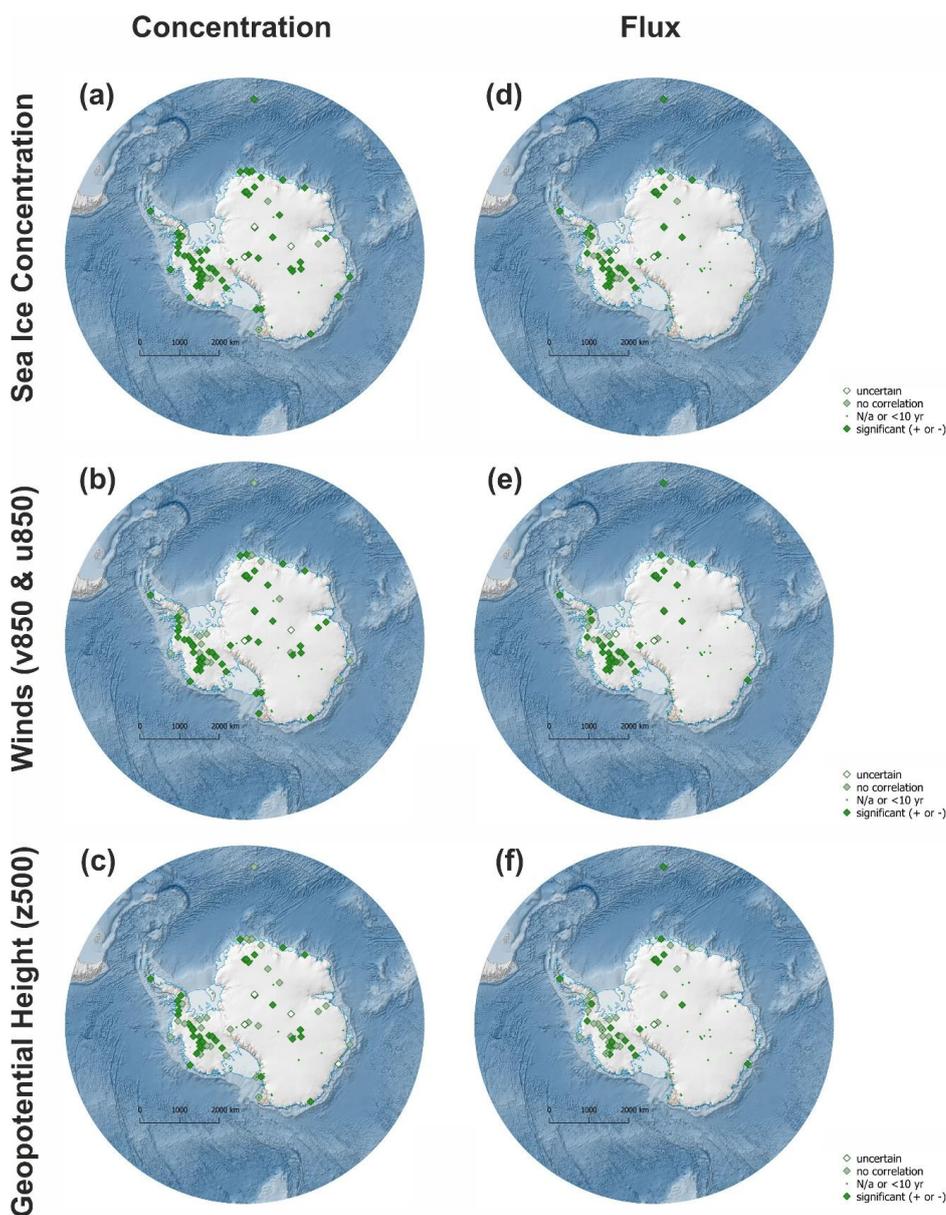


382 **4.4. Excess Sulphate (concentration and flux)**

383 A total of 68 (out of 81) $\text{xs} [\text{SO}_4^{2-}]$ records are correlated with SIC, with five additional records marked as
384 uncertain when using concentration (Table 3). This number drops to 42 (out of 59) when using the flux, with
385 two additional records marked as uncertain. A smaller proportion of records (71 % compared with 84 %)
386 correlated with SIC when using flux.

387 A total of 56 (out of 81) $\text{xs} [\text{SO}_4^{2-}]$ records are correlated winds (v850 and u850), with three additional records
388 marked as uncertain. The number drops to 40 (out of 59) records when using the $\text{xs} \text{SO}_4^{2-}$ flux, with three
389 additional records marked as uncertain. A smaller proportion of records (68% compared with 69 %) correlated
390 with winds when using flux.

391 A total of 40 (out of 81) $\text{xs} [\text{SO}_4^{2-}]$ concentration records are correlated with geopotential height, with an
392 additional six records marked as uncertain. The number drops to 23 (out of 59) records when using the $\text{xs} \text{SO}_4^{2-}$
393 flux, with three additional records marked as uncertain. A smaller proportion of records (39 % compared with
394 49 %) correlated with atmospheric circulation when using flux.



395

396 **Figure 5** – Geographical distribution of xs [SO₄²⁻] records (left column) which exhibit a statistically significant
397 (p>0.05) correlation with (a) SIC, (b) winds (v850 and u850) and (c) geopotential height (z500). Compared with
398 the geographical distribution of xs SO₄²⁻ flux record (right column) which exhibit a statistically significant
399 (p>0.05) correlation with (d) SIC, (e) winds (v850 and u850) and (f) geopotential height (z500). Green
400 diamonds are locations with a significant correlation either positive or negative; grey diamonds are sites with no
401 correlation, open diamonds are uncertain. Dots indicate ice core locations that are in the database but are not
402 less than 10 years in length (or overlap with the instrumental period) or sites which failed to generate any
403 correlations with parameters tested.

404



405 **5. Discussion**

406 **5.1. Which records are suitable for reconstructing SIC, winds and atmospheric circulation?**

407 Our findings reveal that [Na] provides the highest number (69) of records that exhibit a significant correlation
408 with SIC. Only fractionally higher than the number of xs [SO₄²⁻] records (68) and SO₄ (60). Thus, all three
409 records have the potential to capture changes in sea ice conditions. The full list of which sites exhibit positive
410 correlations with each parameter is shown in Supplementary Figure S2.

411 [Na] also provides the highest number of correlations with geopotential height (47) and wind (63). However,
412 proportionally Na flux has the highest number of correlations with geopotential height and winds. While less
413 than 49% of the [SO₄²⁻] and xs [SO₄²⁻] data exhibit relationships with geopotential height, a much higher
414 percentage (64-69 %) display correlations with winds. This suggests that there is greater potential for using
415 [SO₄²⁻] and xs [SO₄²⁻] for reconstructing winds and SIC than geopotential heights. Removing the sea-salt
416 component of [SO₄²⁻] to produce xs [SO₄²⁻] improves the relationship with SIC, geopotential height and winds.

417 Most of the records from West Antarctica and the Antarctic Peninsula (both [Na⁺] and [SO₄²⁻]) exhibit
418 correlations with SIC, geopotential height and winds. This reflects the dominance of marine air-mass incursions
419 in this region (Suzuki et al., 2013), transporting sea salt aerosols from the sea ice zone to the ice core sites. In
420 East Antarctica, the high elevation of the ice sheet (>3000 m) acts as a barrier to marine air-mass transport.
421 However, this study corroborates previous studies (e.g., (Winski et al., 2021)) suggesting that [Na⁺] and [SO₄²⁻]
422 concentrations from ice cores in the East Antarctic plateau are significantly correlated with SIC and atmospheric
423 circulation.

424 Converting the records to flux drastically reduces the geographical coverage. In most cases this is due to the lack
425 of available snow accumulation records from central Antarctica to convert to flux. However, our study
426 demonstrates that converting [Na⁺] to flux increases the relative proportion of records that exhibit a significant
427 correlation with SIC, geopotential height and winds. The opposite is true for [SO₄²⁻] and xs [SO₄²⁻], which
428 results in a lower proportion of records correlating with SIC after converting to flux. This may suggest a
429 dominance of wet deposition of [Na⁺] and dry deposition of [SO₄²⁻]. However, a detailed evaluation of the
430 relationships between ion concentration and snow accumulation is needed to address this fully.

431 Overall, [Na] provides the most records which exhibit significant correlations across all three parameters (179),
432 followed by xs [SO₄²⁻] (164) and [SO₄²⁻] (152).

433 **5.2. Potential limitations**

434 There are limitations to this assessment, which is intended as a first pass filter to highlight the potential future
435 use of the data. In particular, the numbers only relate to records that span or have at least 10-years of data that
436 overlap with the instrumental period. This is defined as the period from 1979-2019 and accounts for 88% of the
437 records (438 out of 499 records submitted). Thus, relationships may exist for shorter records or records drilled
438 prior to 1979, however, it is not possible to verify this under the defined criteria. Another caveat is that
439 correlations have only been conducted with a single sea ice (NSIDC) and reanalysis (ERA-5) product, and
440 results may vary with different datasets. Results may also be impacted by the different timespans used. For
441 example, it was not possible to select the same reference period to run all correlations, because record lengths
442 and top ages (date the core was drilled) vary considerably. Thus, the assumed stationarity in the source and
443 transport routes may not be appropriate.

444 We also note that almost 8% of the records have been classified as “uncertain”. In some cases, significant
445 correlations were evident in the plots, but they were difficult to explain (Fig. 2c). For example, Law Dome
446 generates several regions of significant correlations across multiple sectors, however not in the ocean adjacent to
447 the site. This may indicate long-term transport or the influence of large-scale atmospheric circulation (e.g.,
448 SAM, Indian Ocean Dipole, Atlantic Multidecadal Oscillation). However, in this first pass filter we only
449 included sites where a clear mechanism was evident.

450

451 **6. Data availability**



452 This data descriptor presents version 1.0.0 of the CLIVASH2k Antarctic ice core chemistry database PAGES
453 CLIVASH2k database (Thomas et al., 2022). The database can be accessed via the UK Polar Data Centre.
454 NERC EDS UK Polar Data Centre. <https://doi.org/10.5285/9E0ED16E-F2AB-4372-8DF3-FDE7E388C9A7>

455

456 7. Conclusions.

457 The CLIVASH2k database is the first attempt to compile an Antarctic continental-scale database of chemical
458 records in ice cores spanning the past 2000 years. This study is the first phase of the project, the goal of which
459 was to compile and publish the records. In this study we have provided all available $[\text{Na}^+]$ and $[\text{SO}_4^{2-}]$ records
460 submitted by the community. The records are all available as annual averages, included as both concentration
461 and flux (if available). An additional parameter, $x_s [\text{SO}_4^{2-}]$ has also been calculated where possible.

462 To facilitate future data interpretation, we have run spatial correlations for all the records. The aim of this
463 analysis is to identify sites which exhibit a statistically significant relationship with sea ice concentration (SIC)
464 and atmospheric circulation (500-hPa geopotential heights) or winds (v850 and u850). This is intended as a first
465 filter to identify potential records that could be used in future proxy reconstructions.

466 This first pass filter demonstrates that when considering the species separately, 335 individual records exhibit
467 statistically significant correlations with SIC that have been verified by a team of experts. A recent compilation
468 of available ice core derived sea ice reconstructions, based on a range of proxy data, identified only 17
469 individual sites which have been used to reconstruct sea ice (Thomas et al., 2019). Thus, this data compilation
470 represents a significant improvement on existing published or available data.

471 For researchers interested in reconstructing winds or atmospheric circulation the CLIVASH2k database contains
472 a total of 300 records that are significantly correlated with the wind fields (v850 and u850) and 217 records that
473 are significantly correlated with geopotential height (500 hPa). The Na flux exhibits the greatest proportion of
474 records that correlate with sea ice, atmospheric circulation, and winds. This analysis suggests that Na^+ flux may
475 be the best proxy for reconstructing all three parameters.

476 Future work will focus on using this database to:

- 477 1) Investigate the deposition of $[\text{Na}^+]$ and $[\text{SO}_4^{2-}]$ over decadal to centennial timescales.
- 478 2) Provide a reconstruction of sea ice or atmospheric circulation spanning the past 2000 years.
- 479 3) Evaluate the skill of chemical transport models to capture observed deposition of $[\text{Na}^+]$ and $[\text{SO}_4^{2-}]$.
- 480 4) Combine the information in this new database with the database of snow accumulation (Thomas et al.,
481 2017) and isotopic content (Stenni et al., 2017) to obtain a comprehensive view of Antarctic climate
482 variations over the past 2000 years.

483 This is not an exhaustive list, and we encourage the community to engage with the CLIVASH2k working group
484 and make use of the database.

485

486 Author contributions

487 ET and HG conceived the idea. ET & DV initiated the data call and coordinated the project. ET wrote the paper
488 with contributions from the core writing group. The core writing group (DV, ACFK, DE, HG, DW, VHLW,
489 DD, DU, TV), contributed to the paper writing and discussions. The data interpretation team (ET, DV, ACFK,
490 DW, VHLW, DD, NC, DU, TV, DT, MMG, MS) quality checked the data, evaluated the age-scales, and
491 interpreted the spatial correlation plots. NANB, AH, CML, JRM, YM, KT, HM, YN, FS, JCS, MS, RT, SW,
492 CX, JY, TVK, AAE, LPG and EMT all provided unpublished data. DE wrote the code for the data interpretation
493 plots. DV & LT compiled the figures. All authors read and commented on the manuscript.

494 The following researchers contributed published data to this database. Yoshiyuki Fujii, Lenneke Jong, Elisabeth
495 Isaksson, Filipe G. L. Lindau, Andrew Moy, Rachael Rhodes. We thank the many other researchers who have
496 already made their data available on public data repositories.

497 Competing interests

498 The authors declare no competing conflict of interest.



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