RC4USCoast: A river chemistry dataset for regional ocean model applications in the U.S. East, Gulf of Mexico, and West Coasts

Fabian A. Gomez[1,2], Sang-Ki Lee[2], Charles A. Stock[3], Andrew C. Ross[3], Laure Resplandy[4], Samantha A. Siedlecki[5], Filippos Tagklis[2,6], Joseph E. Salisbury[7]

[1]Northern Gulf Institute, Mississippi State University, Starkville, Mississippi, USA.

[2]NOAA Atlantic Oceanographic and Meteorological Laboratory, Miami, Florida, USA.

[3]NOAA Geophysical Fluid Dynamics Laboratory, Princeton, New Jersey, USA

[4]Department of Geosciences, High Meadows Environmental Institute, Princeton University, Princeton, New Jersey, USA

[5]Department of Marine Sciences, University of Connecticut, Groton, Connecticut, USA

[6]Cooperative Institute for Marine and Atmospheric Studies, University of Miami, Miami, Florida, USA.

[7]Ocean Process Analysis Laboratory, University of New Hampshire, Durham, New Hampshire, USA

Correspondence to: Fabian A. Gomez (fabian.gomez@noaa.gov)

Abstract. A historical dataset of river chemistry and discharge is presented for 140 monitoring sites along the United States East Coast, the Gulf of Mexico, and the West Coast from 1950 to 2022. The dataset, referred to here as River Chemistry for the U.S. Coast (RC4USCoast), is mostly derived from the Water Quality Database of the U.S. Geological Survey (USGS), but also includes river discharge from the USGS's Surface-Water Monthly Statistics for the Nation and the U.S. Army Corps of Engineers. RC4USCoast provides monthly time series as well as long-term averaged monthly climatological patterns for twenty-one variables including alkalinity and dissolved inorganic carbon concentration. It is mainly intended as a data product for regional ocean biogeochemical models and carbonate chemistry studies in the U.S. coastal regions. Here we present the method to derive RC4USCoast and briefly describe the river's carbonate chemistry patterns.

1 Introduction

Riverine fluxes of water, nutrients, alkalinity, and carbon exert a significant impact on the coastal ocean margins, modulating patterns in primary production, dissolved oxygen, calcium carbonate saturation, bottom acidification, and air-sea carbon fluxes (e.g., Rabouille et al., 2008; Cai et al., 2013; Siedlecki et al. 2017; Moore-Maley et al., 2018; Xie et al., 2020; Liu et al., 2021). During the last decade or so, there has been an increasing interest in better understanding and quantifying the influence of river inputs on the coastal ecosystems of the United States. This is reflected in a growing number of ocean biogeochemical (BGC) modeling studies addressing river-induced ocean patterns (e.g., Fennel et al, 2011; 2013; Laurent et al., 2017; Siedlecki et al.

2017; 2021; Hood et al., 2021; Gomez et al., 2021). Ocean BGC models need realistic inputs of river-water properties to properly simulate coastal ecosystem responses to river runoff, but the availability of these inputs is usually limited (e.g., Kearney et al., 2021). A few existing data products contain estimates of riverine carbon and/or nutrients based on empirical or dynamic river export models (e.g., Mayorga et al., 2010; Li et al., 2017; 2019; Lacroix et al., 2021; Regnier et al., 2022). These products were mainly developed for global budget analysis, and consequently they often lack sufficient spatial resolution to allow the study of ecosystem dynamics at a regional scale or have significant regional biases. Motivated by the necessity of high-resolution river chemistry data for regional ocean BGC models, here we present the River Chemistry for the U.S. Coast (RC4USCoast) database, a compilation of historical river chemistry and discharge records derived from the U.S. Geological Survey (USGS).

2 Dataset

The RC4USCoast database contains historical river chemistry records from 140 USGS monitoring stations retrieved from the Water Quality Database of the National Water Information System (Alexander et al., 1998; https://nwis.waterdata.usgs.gov/usa/nwis/qwdata; last access: March 2023). We use a set of stations similar to those used in Stets and Striegl (2012), who selected stations based on the availability of water quality records and proximity to river mouths. These monitoring stations correspond to 52 rivers in the US East Coast, 53 rivers in the Gulf of Mexico, and 35 rivers in the US West Coast (Fig. 1; Table S1 in the Supplement). It is worth noting that Stets and Striegl (2012) reported average inorganic and organic carbon flux (g C yr^{-1}) and yield (g C m^{-2} yr^{-1}) for the selected USGS stations, but they did not provide a dataset with the riverine concentration of carbon. Therefore, RC4USCoast advances providing integrated information on river DIC and alkalinity concentration (Sect. 2.1) and, where available, additional inorganic and organic nutrients relevant for coastal water quality (Sect. 2.2) for those stations.

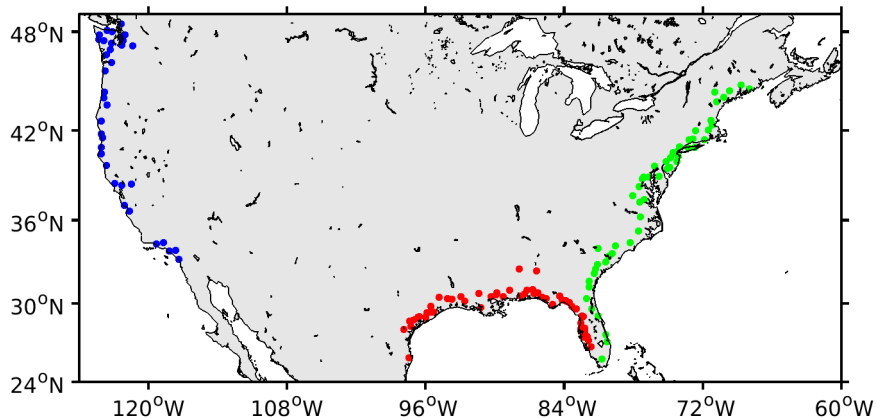



Figure 1. USGS stations used to derive river chemistry patterns. Green, red, and blue dots correspond to rivers discharging to the East, Gulf of Mexico, and West Coast, respectively.

2.1 Carbonate chemistry

RC4USCoast includes a river carbonate chemistry dataset with monthly series and climatological data for alkalinity, pH field, pH laboratory, DIC, and dissolved organic carbon (DOC) (Table 1). To this effect, we processed more than 64,000 records of

55   calcium carbonate ($CaCO_3$) and bicarbonate (HCO_3), 60,000 pH field and laboratory records, and 9,000 DOC records. Due to the substantially smaller number of DIC measurements (~1,000) compared to those of alkalinity and pH, we derived DIC from alkalinity, pH, and water temperature using the CO2SYS program for CO2 System Calculations (van Heuven et al., 2011). Following Stets and Striegl (2012), we assumed that (*i*) particulate inorganic carbon is small; thus, filtered and unfiltered measurements of alkalinity are nearly the same, and (*ii*) organic alkalinity represents a small fraction of total alkalinity. A

60   comparison between filtered and unfiltered measurements of alkalinity does not show significant differences (Fig. 2a); thus, biases associated with the first assumption are negligible. The second assumption is required because including the organic alkalinity fraction in the total alkalinity term used to derive DIC lead to some DIC overestimation. This could be a problem in low alkalinity rivers with high concentration of organic matter, as the latter contains anionic functional groups that can contribute to alkalinity (Hunt et al., 2011). Stets and Striegl (2012) discussed this issue further and showed that organic

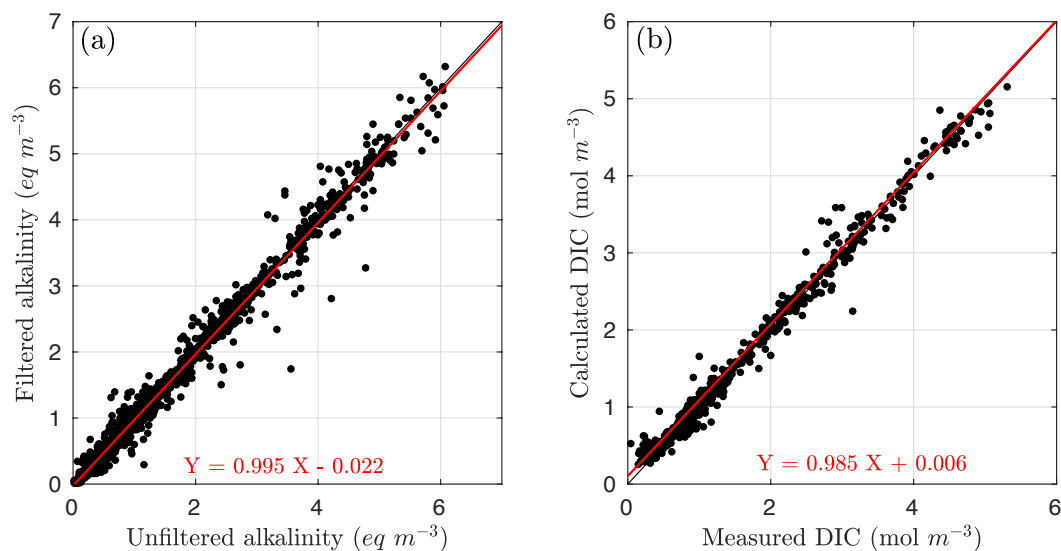
65   alkalinity usually represents less than 10% of total alkalinity in U.S. rivers, not producing important biases in the DIC calculations. Consistently, a comparison between measured DIC and the calculated DIC reveals a good agreement, with no evident bias in the residuals of the least square model (Fig. 2b).

70   **Figure 2.** Data comparison: (a) Filtered vs. unfiltered alkalinity; (b) measured vs. calculated DIC. Calculated DIC was derived from alkalinity, pH, and temperature measurements.

Table 1. Carbon system variables in the RC4USCoast dataset.

Variable[a]	Units	USGS parameter code	Description	Original USGS units	Water chemistry measurements
Alkalinity (alk)	meq m^{-3}	00410	Acid neutralizing capacity, unfiltered, fixed endpoint titration, field	mg CaCO3 liter^{-1}	20,636
		00419	Acid neutralizing capacity, unfiltered, inflection-point titration, field	mg CaCO3 liter^{-1}	393
		29801	Alkalinity, filtered, fixed endpoint titration, laboratory	mg CaCO3 liter^{-1}	3,340
		39036	Alkalinity, filtered, fixed endpoint titration, field	mg CaCO3 liter^{-1}	587
		39086	Alkalinity, filtered, inflection-point titration, field	mg CaCO3 liter^{-1}	7,325
		90410	Acid neutralizing capacity, unfiltered, fixed endpoint titration, laboratory	mg CaCO3 liter^{-1}	8,895
		00440	Bicarbonate, unfiltered, fixed endpoint titration, field	mg HCO3 liter^{-1}	16,168
		00453	Bicarbonate, filtered, fixed endpoint titration, field	mg HCO3 liter^{-1}	7,193
pH field (phf)	standard units	00400	pH, unfiltered, field	standard units	45,866
pH lab (phl)	standard units	00403	pH, unfiltered, laboratory	standard units	14,452
DOC (doc)	mmol C m^{-3}	00681	Organic carbon, filtered	mg C liter^{-1}	9,189
DIC measured (dicm)	mmol C m^{-3}	00691	Inorganic carbon, filtered	mg C liter^{-1}	997
DIC calculated (dic)	mmol C m^{-3}		DIC derived from alkalinity, pH, and temperature		

(a) The abbreviated names of the variables indicated in parenthesis (in first column) are used in the NetCDF files.

75

2.2 Other chemistry variables

The RC4USCoast database also contains a set of variables that describe the runoff of nitrogen, phosphorus, and silica (Table 2), including monthly time series of nitrate (NO_3), nitrate plus nitrite (NO_3 plus NO_2), ammonia (NH_4), organic nitrogen plus ammonia (orgN), dissolved organic nitrogen (DON), total nitrogen (TN), phosphate (PO_4), total phosphorus (TP), and silicon dioxide (SiO_2). For orgN, TN, and TP, we generated two independent datasets for unfiltered and filtered water samples (the former containing both dissolved and particulate material, and the latter only dissolved material). For NO_3, NO_3 plus NO_2, NH_4, and PO_4 we considered the USGS parameters for filtered water samples. In addition to these inorganic and organic nutrients, we also included dissolved oxygen (DO) and water temperature in the database.

80

85   **Table 2.** Additional variables in the RC4USCoast dataset.

Variable[a]	Units	USGS parameter code	Description	Original USGS units	Water chemistry measurements
NO_3 (no3)	mmol N m^{-3}	00618	Nitrate, filtered	mg N liter^{-1}	25,158
		71851	Nitrate, filtered	mg NO3 liter^{-1}	25,012
NO_x (nox)	mmol N m^{-3}	00631	Nitrate plus nitrite, filtered	mg N liter-1	21,879
NH_4 (nh4)	mmol N m^{-3}	71846	Ammonia (NH3 + NH4+), filtered	mg NH4 liter^{-1}	21,609
		00608	Ammonia (NH3 + NH4+), filtered	mg N liter^{-1}	21,360
Organic nitrogen unfiltered (orgNu)	mmol N m^{-3}	00625	Organic nitrogen plus ammonia, unfiltered	mg N liter^{-1}	22,147
Organic nitrogen filtered (orgNf)	mmol N m^{-3}	00623	Organic nitrogen plus ammonia, filtered	mg N liter^{-1}	11,904
DON (don)	mmol N m^{-3}	00607	Dissolved organic nitrogen, filtered	mg N liter^{-1}	11,492
TN unfiltered (tnu)	mmol N m^{-3}	00600	Total nitrogen [inorganic + organic nitrogen], unfiltered	mg N liter^{-1}	24,528
TN filtered (tnf)	mmol N m^{-3}	00602	Total nitrogen [inorganic + organic nitrogen], filtered	mg N liter^{-1}	11,963
PO_4 (po4)	mmol P m^{-3}	00660	Orthophosphate, filtered	mg PO4 liter^{-1}	21,829
		00671	Orthophosphate, filtered	mg P liter^{-1}	20,298
TP unfiltered (tnu)	mmol P m^{-3}	00665	Total phosphorous [organic + inorganic phosphorous], unfiltered	mg P liter^{-1}	28,100
TP filtered (tpf)	mmol P m^{-3}	00666	Total phosphorous [organic + inorganic phosphorous], filtered	mg P liter^{-1}	20,739
Silica (sio2)	mmol Si m^{-3}	00955	Silica, filtered	mg SiO2 liter^{-1}	32,837
Dissolved oxygen (do)	mmol O^2 m^{-3}	00300	Dissolved oxygen, water, unfiltered	mg O2 liter^{-1}	35,615
Temperature (temp)	°C	00010	Water temperature	°C	52,357
Discharge (disc)	m^3 s^{-1}	00060	Mean discharge[b]	ft^3 s^{-1}	
	m^3 s^{-1}	00061	Instantaneous discharge[c]	ft^3 s^{-1}	

(a) The abbreviated names of the variables indicated in parenthesis (in first column) are used in the NetCDF files.

(b) Averaged discharge from the USGS Surface-Water Monthly Statistics was used for all rivers excepting the Mississippi-Atchafalaya (U.S. Army Corps of Engineers) and those listed in c

90   (c) Instantaneous discharge from the USGS Water Quality Database was used for the Charles, James, Weeki Washee, and Rio Grande rivers.

2.3 River discharge

95 To provide a longer set of river discharge records than those available in the USGS Water Quality Database, we used monthly average data from the USGS Surface-Water Monthly Statistics for the Nation database (https://waterdata.usgs.gov/nwis/monthly; last access March 2023). Similarly, for the Mississippi and Atchafalaya rivers, we used records from the U.S. Army Corps of Engineers (USACE; https://rivergages.mvr.usace.army.mil; last access; March 2023). Specifically, we retrieved the Mississippi discharge at the USACE's station 01100 (Tarbert Landing), and the

100 Atchafalaya discharge at station 03045 (Simmesport). For a few rivers (Charles, James, Weeki Washee, and Rio Grande) where monthly discharge was not available in the USGS Surface-Monthly Statistics database or the USACE records, we used discharge from the USGS Water Quality Database.

2.4 Database generation

Information for the selected river stations includes the RC4USCoast river ID, the original USGS site ID, the USGS site's

105 longitude and latitude, and an approximate longitude and latitude for the river mouth (Fig. S1 in the Supplement). A few rivers flow to other larger rivers, as described in Table S1. The assigned mouth location in those cases corresponds to the mouth of the major stream discharging to the ocean. For example, the dataset contains the Alabama and Tombigbee rivers, which converge to the Mobile River, so the associated river mouth for those two rivers is the Mobile mouth (30.7°N and 88.0°W).

To the extent it was possible given data availability, we calculated monthly times series for all variables and all river sites over

110 the period 1950–2022. Temporal data gaps were kept unfilled. In the Water Quality Database, river properties are characterized by a set of parameters, each associated with a specific measurement type. As indicated in Tables 1 and 2, we used eight parameters to derive alkalinity, two parameters to derive NO_3, NH_4 and PO_4, and one parameter for the remaining variables: pH field, pH laboratory, DOC, measured DIC, NO_3 plus NO_2, SiO_2, DO, temperature, and the filtered and unfiltered concentration of orgN, TN, and TP. Using more than one parameter for alkalinity, NO_3, NH_4 and PO_4 was a reasonable option

115 to improve the spatiotemporal representation of the patterns, as concentration differences between parameters, during overlapping periods, were minor and did not reveal evident biases (Figs. S2 and S3 in the Supplement). Conversion factors were applied to present alkalinity in milliequivalent m^{-3} (meq m^{-3}), the carbon-based variables (DIC, DOC) in mmol C m^{-3}, the nitrogen-based variables (NO_3, NO_3 plus NO_2, NH_4, orgN, TN) in mmol of N m^{-3}, the phosphorous-based variables (PO_4 and TP) in mmol of P m^{-3}, silica in mmol of SiO_2 m^{-3}, and dissolved oxygen in mmol of O_2 m^{-3}. To ensure data quality, outliers,

120 defined here as river chemistry values above and below 3.5 standard deviations from the median were removed. Maximum alkalinity (and calculated DIC) values were limited to 8,000 meq (mmol) m^{-3}. pH records below 3.5 or above 10 units were discarded. Additionally, an upper threshold of 3.5 was used for the DIC to alkalinity ratio (DIC:Alk ratio), based on values reported by Moore-Maley et al. (2018). Calculated DIC records linked to DIC:Alk ratios greater than 3.5 were then removed.

This was a very minor fraction of the total monthly DIC records (3.7%), mainly associated with low alkalinity's values in the Toms, Satilla, St. Marys, and Blackwater rivers.

Except for river discharge, which had average temporal coverage of 85%, the USGS time series had significant data gaps because the parameter's monitoring had a limited number of years, and/or the parameter's measurements were not performed at a regular frequency. Figure 3 displays the number of records (data density) in the monthly time series for each site. We omitted measured DIC, which was present at a low number in only 10% of the stations. The density maps reveal large differences among rivers and variables. Monitoring stations with the most complete chemistry records were linked to rivers flowing to the Mid Atlantic Bight, the Mississippi and Atchafalaya, and a limited number of major rivers on the West Coast and Texas coast. The greatest data density was for water temperature, pH field, and alkalinity, with a median of 176, 170, and 149 records (over the 140 sites), respectively, followed by dissolved oxygen, silica, and calculated DIC, with a median of 138, 134, and 124 records, respectively. At the other end, variables with low data density were DON and DON, with a median of 21 and 14 records, respectively.

To complement the time series and provide a ready to use dataset for ocean biogeochemical model applications with no data gaps, we generated monthly climatologies using all data during 1950–2022. We also generated climatologies for the 1950–1989 and 1990–2022 periods, as a way to represent temporal variation in the climatological pattern. We considered those multidecadal periods, as the temporal coverage in the river chemistry dataset did not resolve well decadal variability for all sites. To ensure a minimum number of observations to derive the monthly climatologies, for each variable and station we calculated the number of records per calendar month. If the median value of this record count (over the 12 months) was less than five, or any month had no data, then the monthly climatology was substituted by the long-term annual average.

A brief description of the carbon system variables in the RC4USCoast database (based on alkalinity and calculated DIC) is provided in the following section. Mean patterns for other variables are shown in the Supplement (Fig. S4).

3 Main carbon system patterns

The site-averaged alkalinity concentration ranges from 40 meq m^{-3} (Black water) to 5,605 meq m^{-3} (Santa Clara). The frequency distribution for this variable displays a positive skewness with a median of 662 meq m^{-3} and 42% of the values lower than 500 meq m^{-3} (Fig. 4a). The largest fraction of low alkalinity (<500 meq m^{-3}) rivers is on the East Coast, especially for rivers flowing to the Gulf of Maine and South Atlantic Bight (Fig. 5a). On the other hand, the largest fraction of high alkalinity (>2,000 meq m^{-3}) rivers is in the Gulf of Mexico (Fig. 4a), mainly clustered over the Texas and West Florida coasts (Fig. 5a). Along the West Coast, there is a clear meridional gradient in river alkalinity, with the highest values in Southern California and the lowest in Oregon and Washington (Fig. 5a).

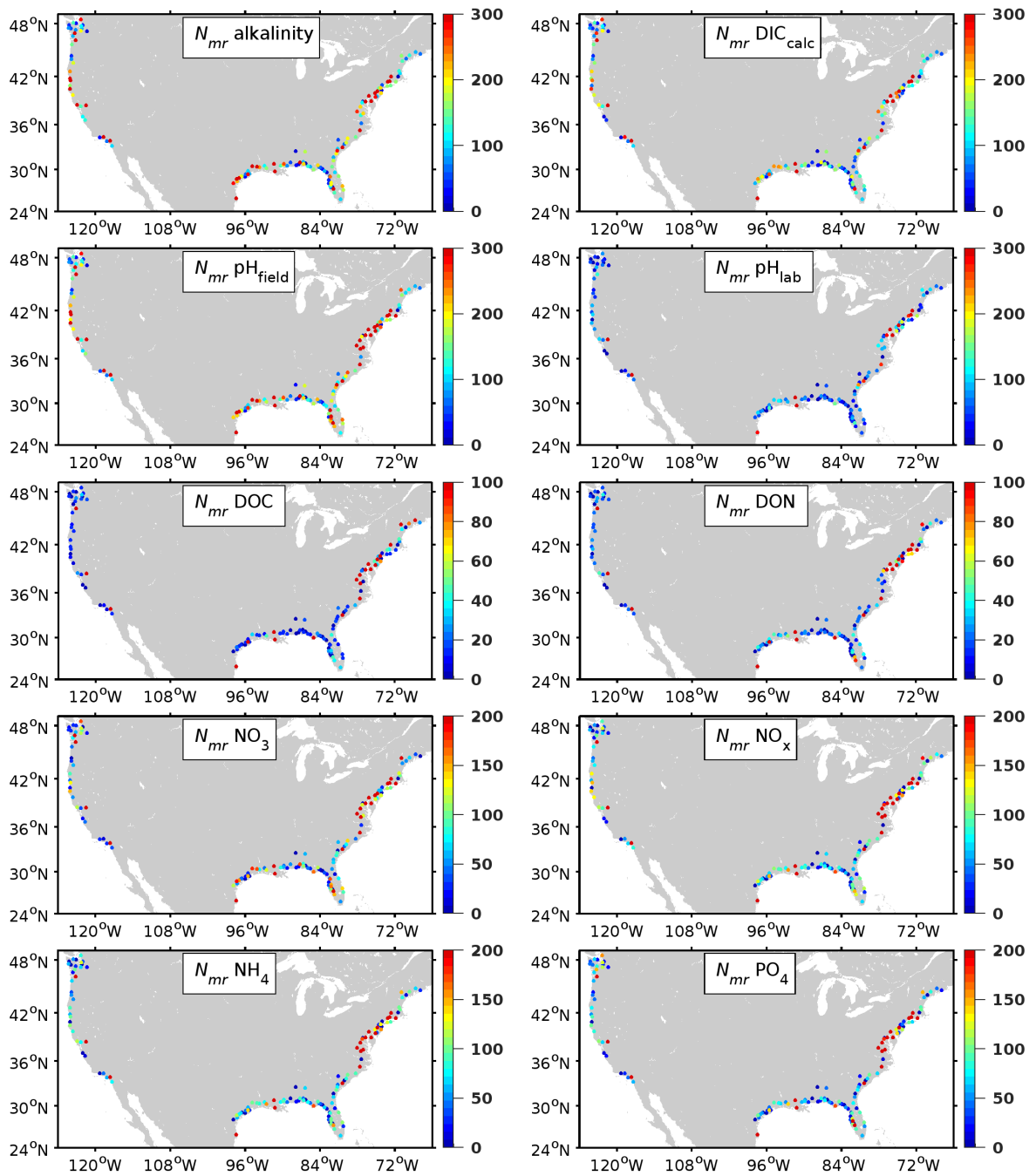



Figure 3. Number of monthly records (N_{mr}) in the dataset time series (1950-2022). The colorbar range may vary between panels. Variable description is in Tables 1 and 2.

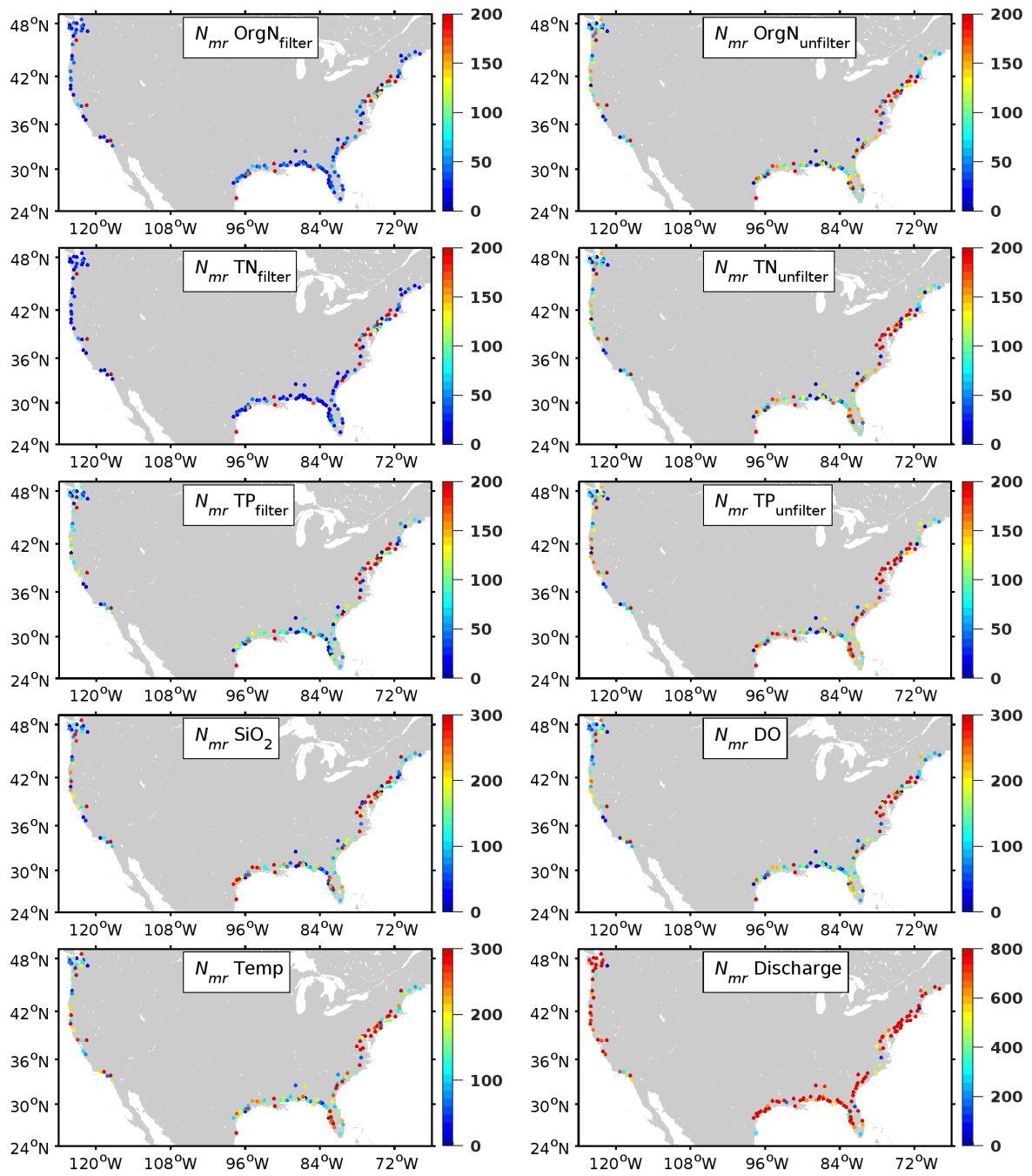



Figure 3 (continued).

160

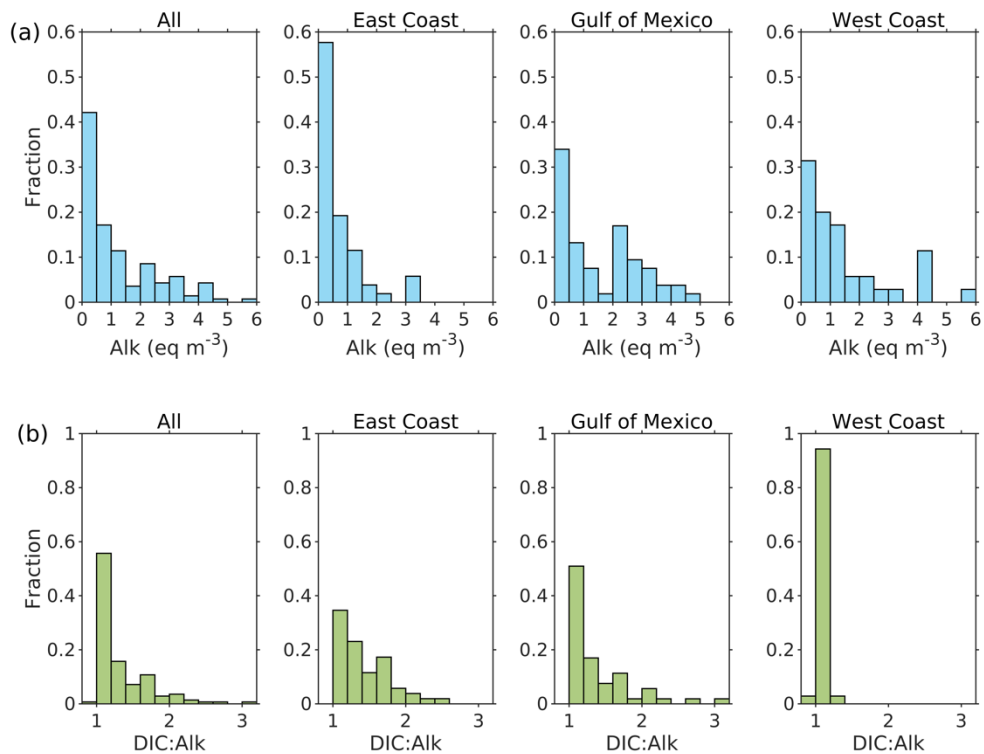



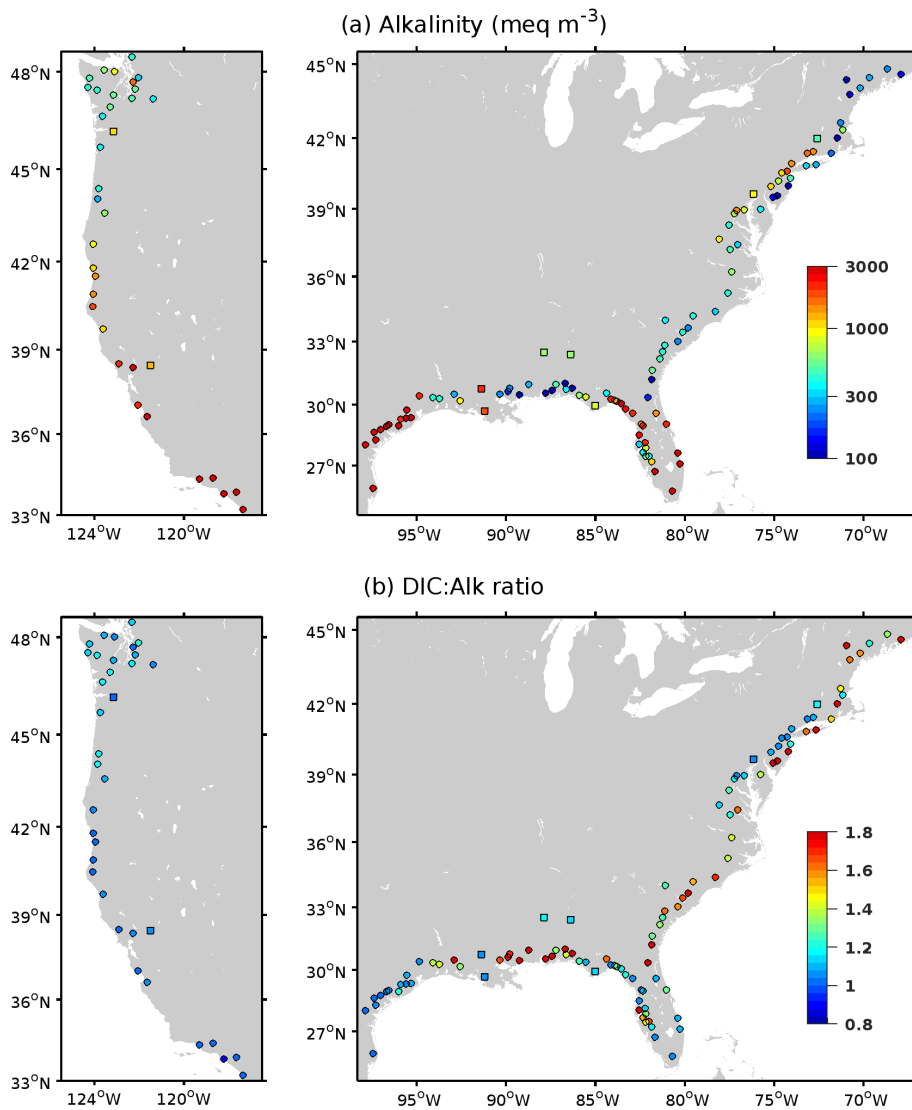
Figure 4. Frequency histogram derived from the long-term site-averaged (a) alkalinity and (b) DIC to alkalinity (DIC:Alk) ratio for all rivers (All), and river discharging at the East Coast, Gulf of Mexico, and West Coast.

165

Figure 5. Long-term mean (colored dots and squares) of the river (a) alkalinity and (b) DIC to alkalinity ratio. Squares (dots) represent river stations with a mean discharge greater (smaller) than 500 m³ s⁻¹. Colorbar in (a) is in logscale.

The average river DIC concentration shows a very similar spatial pattern to the average river alkalinity, as both variables are
170   highly correlated ($r = 0.99$; Fig. S5 in the Supplement). However, DIC tends to be greater than alkalinity, which is reflected in
an average DIC:Alk ratio of 1.33 over the 140 stations. Like alkalinity, the frequency distribution of the site averaged DIC:Alk
ratios has a positively skewed distribution (Fig. 4b), with a median of 1.17, and minimum and maximum values of 0.92 (Los
Angeles) and 3.09 (Shoal), respectively. Rivers with the lowest DIC:Alk ratios are in the West Coast, where DIC is on average
8% greater than alkalinity (Fig. 5b). Large DIC:Alk ratios are mainly associated with low alkalinity rivers, and the opposite is
175   true for high alkalinity rivers. Indeed, the relationship between these two variables has a clear linear pattern for alkalinities
below ~500 meq m^{-3}, where the mean DIC:Alk ratio decreases 0.234 units per every 100 meq alkalinity increase (Fig. 6a).
Moreover, we found that the standard deviation of the DIC:Alk ratio (SD$_{DIC:Alk}$) is inversely linked to the mean alkalinity (Fig.
6b). Most rivers with a mean alkalinity below 200 meq m^{-3} have a SD$_{DIC:Alk}$ greater than 0.4, whereas most rivers with a mean
alkalinity above 1,000 meq m^{-3} have a SD$_{DIC:Alk}$ lower than 0.2.

180

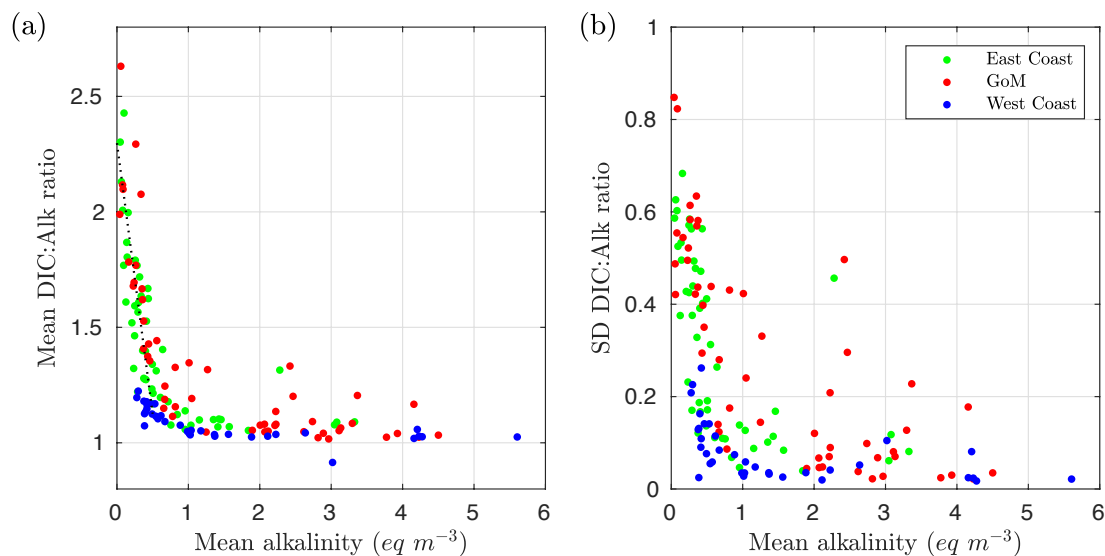



Figure 6. Between river variability in the DIC to alkalinity (DIC:Alk) ratio as a function of alkalinity: (a) mean DIC:Alk ratio
vs. mean alkalinity; b) standard deviation of the DIC:Alk ratio vs. mean alkalinity. Each dot represents one of 140 rivers in the
dataset. Green, red, and blue dots depict the rivers flowing to the East, Gulf of Mexico, and West Coasts, respectively.

185

To further investigate the river carbonate chemistry variability, we examined monthly records for the stations with the largest data density. Those stations are associated with six rivers in the East Coast (Connecticut, Delaware, Schuylkill, Choptank, Susquehanna, and Neuse), two rivers in the Gulf of Mexico (Mississippi and Atchafalaya) and four rivers in the West Coast (Santa Ana, Sacramento, Eel, and Klamath). A strong positive relationship between the monthly alkalinity and DIC records is evident for all of the 12 stations (Fig. 7a). The coefficients of determination (R^2) for the linear regression of DIC against alkalinity average to 0.89, ranging from 0.57 (Neuse) to 0.99 (Eel) (Table S2 in the Supplement). Like the patterns in Figure 6, the monthly records show an inverse relationship between the DIC:Alk ratio and alkalinity (i.e., an increased variability in the DIC: Alk ratio at low alkalinity values, and vice versa at high alkalinity values) (Fig.7b). The Choptank and Neuse rivers, in the lower end of the alkalinity concentration, show the largest dispersion in the DIC:Alk ratio, with values ranging from ~1 to higher than 2.5. In contrast, the high alkalinity Santa Ana River displays a much smaller variability, with the maximum DIC:Alk ratio around 1.1. Seasonal patterns for alkalinity (and DIC) tend to show enhanced values during summer and fall, and minimum values during winter and spring (Fig. 7c), concurrent with low and high discharge periods, respectively. This pattern is consistent with multiple studies conducted in specific river basins suggesting dilution of carbonate chemistry variables during high discharge conditions (e.g., Cai, 2003; Guo et al., 2008; Joesoef et al., 2017). Indeed, a linear relationship between the logarithm of discharge (logDisc) and alkalinity is evident for each of the 12 stations (Fig. 7d). The adjusted linear regression models for these rivers are all significant, with linear regression coefficients ranging from 0.33 (Sacramento) to 0.69 (Eel) (Table S2).

The inverse relationship between logDisc and alkalinity in Figure 7d can be extended to other rivers in the database. Figure 8 shows the regression coefficient (slope) and R^2 for the stations where the regression was significant, explained at least 20% of the alkalinity variance, and include at least 30 observations (78 out of 140 rivers). The sensitivity of river alkalinity to changes in discharge, reflected in the magnitude of the regression coefficient, is greater in the high alkalinity rivers flowing to the Gulf of Mexico, Southern California, and East Florida coasts, and smaller in the low alkalinity rivers flowing to the Northwest and East Coasts (Fig. 8a). This determines a significant negative correlation between the regression coefficient and the site-averaged alkalinity ($r = -0.69$). The R^2 coefficient pattern shows an important spatial variability (Fig. 8b), which is not linked to river alkalinity or discharge. The largest R^2 values (>0.5) characterize rivers flowing to the Northwest Coast, Florida Panhandle, and South Atlantic Bight. Similar patterns were found for the relationship between logDisc and DIC (not shown).

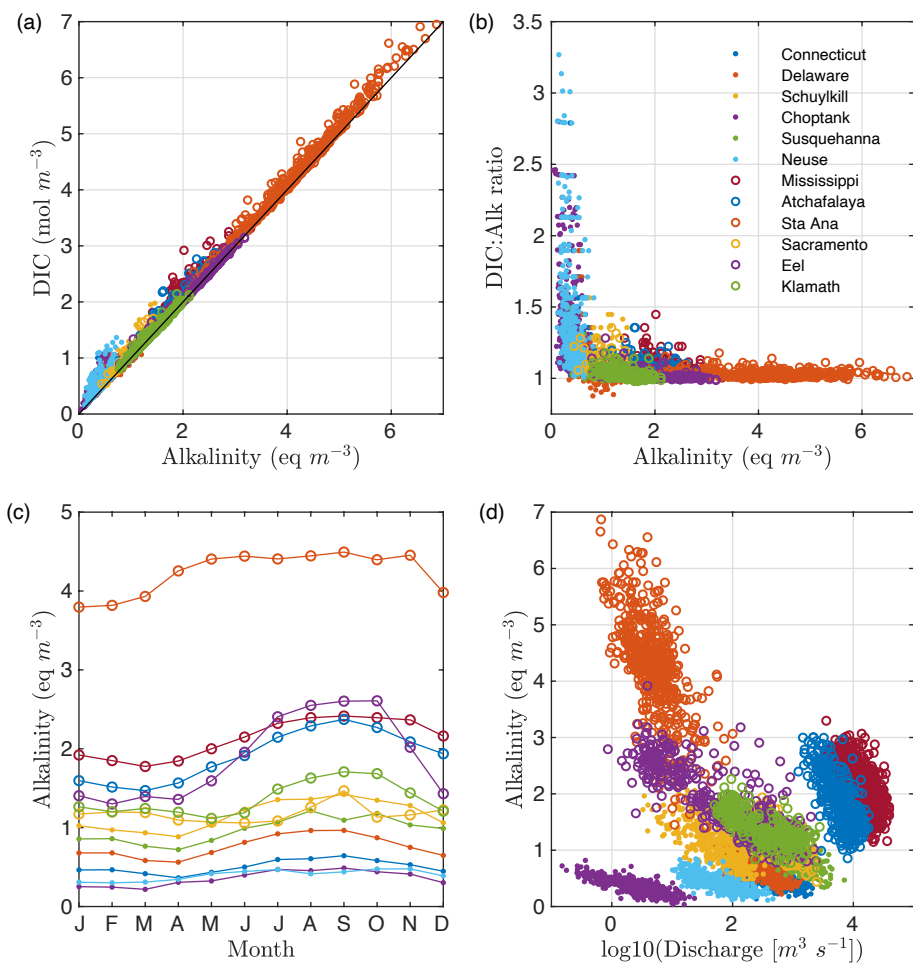



Figure 7. Carbon system patterns for 12 selected rivers: (a) DIC vs. alkalinity, and (b) DIC:Alk ratio vs. alkalinity; (c) monthly climatological patterns of alkalinity; and (d) alkalinity vs. logarithm of discharge. All patterns were derived for the 12 rivers with the largest number of records in the database: Connecticut (ID=11), Delaware (23), Schuylkill (24), Choptank (25), Susquehanna (26), Neuse (35), Mississippi (86), Atchafalaya (88), Santa Ana (107), Sacramento (113), Eel (117), and Klamath (119).

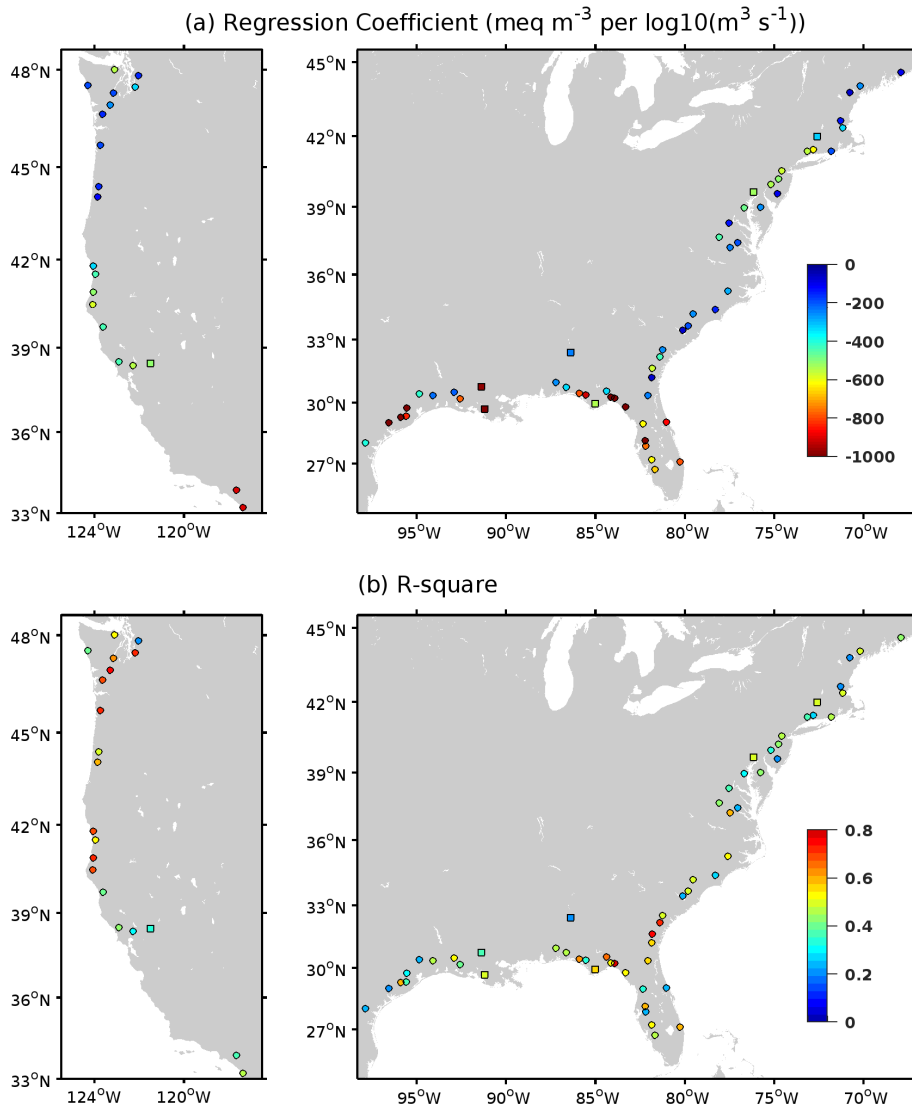



Figure 8. (a) Regression coefficient and (b) coefficient of determination (R^2) for the adjusted linear regressions between alkalinity and the logarithm of discharge (colored dots and squares). Patterns are shown only for the stations that have a significant regression coefficient, an R^2 greater than 0.2, and more than 30 observations (78 out of 140 stations). Squares (dots) represent river stations with a mean discharge greater (smaller) than 500 m^3 s^{-1}.

4 Data availability

The river chemistry data product is available in netCDF format at NOAA/NCEI with a DOI of
230   https://doi.org/10.25921/9jfw-ph50 and NCEI accession number 0260455 (Gomez et al., 2022). For each of the selected river stations, we provide monthly time series and climatologies for each variable. Excel spreadsheets reporting the USGS parameters used to generate the dataset, the station and river mouth locations (latitude/longitude), the number of records in the series, and the first and last year in the series, are also provided in the dataset.

5 Summary and conclusion

235   Retrieving data from the USGS Water Quality database has complexities, such as identifying nearshore sites for coastal studies (USGS contains more than 2,400 sites across the United States, many in inland waters that are not directly relevant to coastal ocean analyses), or integrating water quality parameters to characterize biogeochemical properties (water properties are usually described by more than one USGS parameter). Thus, a user not familiar with the USGS database may require considerable time and effort identifying river sites and parameters. We facilitate this task, providing an integrated river chemistry and
240   discharge dataset for 140 USGS nearshore sites, which contains relevant variables to characterize biogeochemical and water fluxes (land-to-ocean) along the U.S. West, East and Gulf of Mexico coasts. RC4USCoast includes data for alkalinity, pH, nutrients, and novel estimates of river DIC. River mouth location (longitude, latitude) is reported for each USGS sites, which expedites the data integration in coastal biogeochemical studies. The main goal is to fill a gap for river carbonate chemistry products, as necessary inputs for regional model simulations that include ocean biogeochemistry. We also note the utility of
245   this product for skill assessment of hydrologic and riverine chemistry models estimating discharge and nutrient loading patterns resulting from climate and land use activities (e.g., Lee et al., 2019). Patterns in RC4USCoast show distinct regional features for alkalinity and DIC. The average and standard deviation of the DIC:Alk ratio increased in low alkalinity rivers, and both alkalinity and DIC concentration were inversely related to river discharge. Our results revealed a significant spatiotemporal variability in carbonate chemistry, which can play a significant role on coastal biogeochemical dynamics.

250   **Author contributions:** FAG retrieved the USGS dataset and generated the monthly series and climatological data for the selected sites and variables. AR produced the files in netCDF format. All authors contributed to the writing of the paper.

Competing interests: The authors declare that they have no conflict of interest.

Acknowledgements: This article was supported by NOAA's Ocean Acidification Program, Climate Program Office, and Atlantic Oceanographic and Meteorological Laboratory.

255   **Financial support:** The research was conducted under NOAA's awards to the Northern Gulf Institute (NA16OAR4320199).

References

Alexander, R.B., Slack, J.R., Ludtke, A.S., Fitzgerald, K.K. and Schertz, T.L. Data from selected US Geological Survey national stream water quality monitoring networks. *Water Resources Research*, *34*(9), 2401-2405, https://doi.org/10.1029/98WR01530, 1998.

260 Cai, W.-J.: Riverine inorganic carbon flux and rate of biological uptake in the Mississippi River plume, Geophys. Res. Lett., 30, 1032, https://doi.org/10.1029/2002GL016312, 2003.

Fennel, K., Hetland, R., Feng, Y., and DiMarco, S.: A coupled physical-biological model of the Northern Gulf of Mexico shelf: model description, validation and analysis of phytoplankton variability, Biogeosciences, 8, 1881–1899, https://doi.org/10.5194/bg-8-1881-2011, 2011.

265 Fennel, K., Hu, J., Laurent, A., Marta-Almeida, M., and Hetland, R.: Sensitivity of hypoxia predictions for the northern Gulf of Mexico to sediment oxygen consumption and model nesting, J. Geophys. Res.-Ocean, 118, 990–1002, 2013.

Gomez, F. A., Wanninkhof, R., Barbero, L., and Lee, S. K.: Increasing river alkalinity slows ocean acidification in the northern Gulf of Mexico, Geophysical Research Letters, https://doi.org/10.1029/2021GL096521, 2021.

Gomez, F.A., Lee, S.K., Stock, C.A., Ross, A.C., Resplandy, L., Siedlecki, S.A., Tagklis, F., Salisbury, J.E.: RC4USCoast: A
270 river chemistry dataset for regional ocean model application in the U.S. East, Gulf of Mexico, and West Coasts from 1950-01-01 to 2020-12-31 (NCEI Accession 0260455), NOAA National Centers for Environmental Information, Dataset https://doi.org/10.25921/9jfw-ph50, 2022.

Guo, X., Cai, W.J., Zhai, W., Dai, M., Wang, Y. and Chen, B: Seasonal variations in the inorganic carbon system in the Pearl River (Zhujiang) estuary, *Continental Shelf Research*, *28*(12), 1424-1434, https://doi.org/10.1016/j.csr.2007.07.011, 2008.

275 Hood, R. R., Shenk, G. W., Dixon, R. L., Smith, S. M., Ball, W. P., Bash, J. O. et al.: The Chesapeake Bay program modeling system: Overview and recommendations for future development. *Ecological Modelling*, *456*, 109635, https://doi.org/10.1016/j.ecolmodel.2021.109635, 2021

Hunt, C. W., Salisbury, J. E., and Vandemark, D.: Contribution of non-carbonate anions to total alkalinity and overestimation of pCO$_2$ in New England and New Brunswick rivers, Biogeosciences, 8, 3069–3076, https://doi.org/10.5194/bg-8-3069-2011,
280 2011.

Joesoef, A., Kirchman, D. L., Sommerfield, C. K., and Cai, W.-J.: Seasonal variability of the inorganic carbon system in a large coastal plain estuary, Biogeosciences, 14, 4949–4963, https://doi.org/10.5194/bg-14-4949-2017, 2017.

Kaushal, S.S., Likens, G.E., Utz, R.M., Pace, M.L., Grese, M. and Yepsen, M. Increased river alkalinization in the Eastern US. *Environmental science & technology*, *47*(18), pp.10302-10311, 2013.

285   Kearney, K. A., Bograd, S. J., Drenkard, E., Gomez, F. A., Haltuch, M., Hermann, A. J., Jacox, M. G., Kaplan, I. C., Koenigstein, S., Luo, J. Y., and Masi, M.: Using global-scale earth system models for regional fisheries applications, Frontiers in Marine Science, 1121, https://doi.org/10.3389/fmars.2021.773443, 2021.

Mayorga, E., Seitzinger, S. P., Harrison, J. A., Dumont, E., Beusen, A. H., Bouwman, A. F., Fekete, B. M., Kroeze, C., and Van Drecht, G.: Global nutrient export from WaterSheds 2 (NEWS 2): model development and implementation,
290   Environmental Modelling and Software, 25(7), pp.837-853, https://doi.org/10.1016/j.envsoft.2010.01.007, 2010.

Lacroix, F., Ilyina, T., Mathis, M., Laruelle, G. G., and Regnier, P.: Historical increases in land-derived nutrient inputs may alleviate effects of a changing physical climate on the oceanic carbon cycle, Global change biology, 27(21), 5491-5513, https://doi.org/10.1111/gcb.15822, 2021.

Lacroix, F., Ilyina, T., and Hartmann J.: Oceanic CO_2 outgassing and biological production hotspots induced by pre-industrial
295   river loads of nutrients and carbon in a global modelling approach, Biogeosciences, 17, 55–88, https://doi.org/10.5194/bg-17-55-2020, 2020.

Lee, M., Shevliakova, E., Stock, C. A., Malyshev, S., and Milly, P. C.: Prominence of the tropics in the recent rise of global nitrogen pollution, Nature communications, 10(1), 1-11, https://doi.org/10.1038/s41467-019-09468-4, 2019.

Li, M., Peng, C., Wang, M., Xue, W., Zhang, K., Wang, K., Shi, G., and Zhu, Q.: The carbon flux of global rivers: a re-
300   evaluation of amount and spatial patterns, Ecological Indicators, 80, 40-51, https://doi.org/10.1016/j.ecolind.2017.04.049, 2017.

Li, M., Peng, C., Zhou, X., Yang, Y., Guo, Y., Shi, G. and Zhu, Q.: Modeling global riverine DOC flux dynamics from 1951 to 2015, Journal of Advances in Modeling Earth Systems, 11(2), 514-530, https://doi.org/10.1029/2018MS001363, 2019.

Liu, X., Stock, C.A., Dunne, J.P., Lee, M., Shevliakova, E., Malyshev, S. and Milly, P.C.: Simulated Global Coastal Ecosystem
305   Responses to a Half-Century Increase in River Nitrogen Loads. *Geophysical Research Letters*, *48*(17), p.e2021GL094367, https://doi.org/10.1029/2021GL094367, 2021.

Moore-Maley, B. L., Ianson, D., and Allen, S. E.: The sensitivity of estuarine aragonite saturation state and pH to the carbonate chemistry of a freshet-dominated river, Biogeosciences, 15, 3743–3760. https://doi.org/10.5194/bg-15-3743-2018, 2018.

Rabouille, C., Conley, D. J., Dai, M. H., Cai, W. J., Chen, C. T. A., Lansard, B., Green, R., Yin, K., Harrison, P. J., Dagg, M.,
310   and McKee, B.: Comparison of hypoxia among four river-dominated ocean margins: The Changjiang (Yangtze), Mississippi, Pearl, and Rhône rivers, Continental Shelf Research, 28(12), 1527-1537, https://doi.org/10.1016/j.csr.2008.01.020, 2008.

Regnier, P., Resplandy, L., Najjar, R.G., and Ciais, P.: The land-to-ocean loops of the global carbon cycle, Nature, 603, 401–410, https://doi.org/10.1038/s41586-021-04339-9, 2022.

Siedlecki, S. A., Salisbury, J., Gledhill, D. K., Bastidas, C., Meseck, S., McGarry, K., Hunt, C. W., Alexander, M., Lavoie, D., Wang, Z. A., and Scott, J.: Projecting ocean acidification impacts for the Gulf of Maine to 2050: New tools and expectations, Elem Sci Anth, 9(1), p.00062. https://doi.org/10.1525/elementa.2020.00062, 2021.

Siedlecki, S. A., Pilcher, D. J., Hermann, A. J., Coyle, K., and Mathis, J.: The importance of freshwater to spatial variability of aragonite saturation state in the Gulf of Alaska, Journal of Geophysical Research: Oceans, 122(11), 8482-8502, https://doi.org/10.1002/2017JC012791, 2017.

Stets, E. G., and Striegl, R. G.: Carbon export by rivers draining the conterminous United States, Inl. Waters, 2, 177–184, https://doi.org/10.5268/IW-2.4.510, 2012.

Stets, E.G., Kelly, V.J. and Crawford, C.G.: Long-term trends in alkalinity in large rivers of the conterminous US in relation to acidification, agriculture, and hydrologic modification. *Science of the Total Environment*, *488*, pp.280-289, https://doi.org/10.1016/j.scitotenv.2014.04.054, 2014.

van Heuven, S. M. A. C., Pierrot, D., Rae, J. W. B., Lewis E., and Wallace D. W. R.: MATLAB program developed for CO2 system calculations, ORNL/CDIAC-105b Carbon Dioxide Information Analysis Center, Oak Ridge National Laboratory, US Department of Energy, Oak Ridge, Tennessee, 530, 2011.

Xie, Y., Lin, L., Xiao, W., Yu, X., Lan, W., and Huang, B.: Striking seasonal pattern of primary production in the river-dominated ocean margin of the northern South China Sea (NSCS-RiOMar) revealed by new field and remotely sensed data, Progress in Oceanography, 189, p.102470, https://doi.org/10.1016/j.pocean.2020.102470, 2020.