

1 **Integrated ecohydrological hydrometric and stable water isotope data**
2 **of a drought-sensitive mixed land use lowland catchment**

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14 **Abstract**

15 Data from long-term experimental catchments are the foundation of hydrological sciences and
16 are crucial to benchmark process understanding, observe trends and natural cycles, and are
17 prerequisites for testing predictive models. Integrated data sets which capture all compartments
18 of our landscapes are particularly important in times of land use and climate change. Here, we
19 present ecohydrological data measured at multiple spatial scales which allows differentiation of
20 “blue” water fluxes (which maintain streamflow generation and groundwater recharge) and
21 “green” water fluxes (which sustain vegetation growth). There are two particular unique aspects
22 to this data set : a) we measured water stable isotopes in the different landscape compartments
23 (that is in precipitation, surface water, soil, ground- and plant water); and b) we conducted this
24 monitoring during the extreme drought of 2018 in Central Europe. Stable water isotopes are so
25 useful in hydrology as they provide “fingerprints” of the pathways water took when moving
26 through a catchment. Thus, isotopes allow to evaluate the dynamic relationships between water
27 storage changes and fluxes, which is fundamental to understanding how catchments respond to
28 hydroclimate perturbations or abrupt land use conversion. Second, as we provide the data until
29 2020 one can also investigate recovery of water stores and fluxes after extreme droughts. Last
30 but not least: lowland headwaters are often understudied systems despite them providing
31 important ecosystem services such as groundwater and drinking water provision and
32 management for forestry and agriculture.

34 1. Introduction

35

36 Progress in scientific hydrology and provision of an evidence base for sustainable land and water
37 management are only possible due to detailed, long-term observational data collected from
38 experimental watersheds (Hewlett et al., 1969; Robinson et al., 2013). Such experimental
39 “outdoor laboratories” are invaluable scientific resources given the complexity of increasing
40 pressures on water supplies (e.g. Cosgrove and Loucks, 2015), land use change (Neill et al.,
41 2021) and the uncertain effects and non-stationarity of projected climate change (Milly et al.,
42 2015).

43 Ecohydrology adopts an interdisciplinary approach to investigate links between the structure and
44 function of ecological systems and the partitioning, flux and storage of fresh water (Guswa et al.,
45 2020). Recent advances in monitoring and modeling have created manifold opportunities to
46 address urgent ecohydrological questions on the importance of links between processes across
47 the critical zone (CZ) - the dynamic, life-sustaining near-surface of the terrestrial earth that
48 extends between the top of vegetation canopies, through the soil and into groundwater (Grant &
49 Dietrich, 2017). Within the CZ concept, vegetation plays a central and dynamic role in partitioning
50 incoming precipitation into “blue water” fluxes (streamflow generation and groundwater recharge)
51 and “green water” fluxes which maintain vegetation growth (Evaristo et al., 2015).

52 To enhance ecohydrological process understanding in catchment systems, robust, multi-scale
53 integrated data sets are required (Tetzlaff et al., 2021). In this regard, water stable isotopes and
54 other tracers can help identify sources and pathways of water in the landscape and across the
55 CZ to elucidate how different land use affects water partitioning between green and blue water
56 fluxes (Dubbert and Werner, 2019; Tetzlaff et al., 2015). Importantly, water stable isotopes have
57 enhanced the characterization of the celerity of hydrological fluxes in different CZ compartments,
58 as well as quantifying the velocity of water particles and associated mixing relationships in the
59 subsurface (Benettin et al., 2015; Birkel et al., 2011). Evaluating the dynamic relationships
60 between water storage changes and fluxes is fundamental to understanding how catchments
61 respond to hydroclimate perturbations, such as anomalous dry or wet periods, or abrupt land use
62 conversion. This provides a more nuanced and integrated understanding of how key
63 ecohydrological couplings may be at risk during long-term changes in blue and green water
64 partitioning resulting from climate and land use change (Orth and Destouni, 2018). Such
65 integrated understanding is important in the context of projected increases in air temperature,
66 aridity, and in precipitation patterns, which may cause more variability in water availability
67 threatening the sustainability of important ecosystem services (Okruszko et al., 2011). As an
68 increase in drought frequency and severity is expected across Europe as the 21st century
69 progresses, the development of effective and evidence-based amelioration measures to underpin
70 sustainable and integrated land and water management policies for changing climatic conditions
71 is urgently needed (Samaniego et al., 2018).

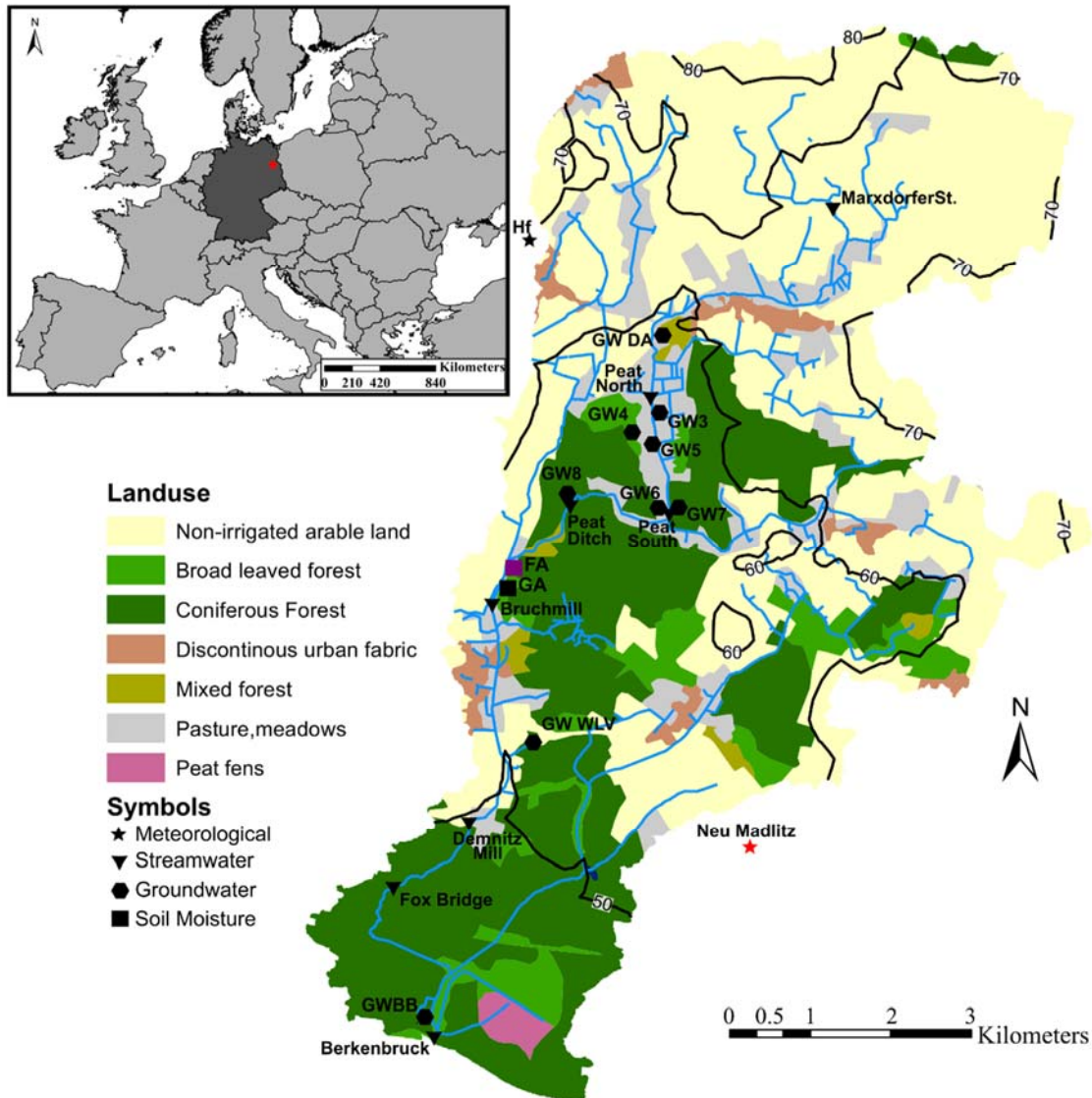
72 Consequently, integrated ecohydrological and stable isotope data sets targeted at understanding
73 the effects of different types of environmental change have outstanding potential, not least
74 because interdisciplinary environmental research tends to give unanticipated insights (Burt,
75 1994). Such integrated data streams allow identification and quantification of the linkages
76 between rainfall, soil moisture, groundwater and runoff generation, facilitating deeper
77 understanding of flood and drought risk in different types of landscapes and under different land
78 use management (Huntingford et al., 2014).

79 Water resources in the extensive, glacially formed, lowland landscape of northern Europe,
80 including the North German Plain sustain food production (Gutzler et al., 2015; Barkmann et al.,
81 2017) and water supplies to large cities like Berlin. Interestingly, such lowland catchments are still
82 relatively understudied compared to more upland headwater landscapes with stronger
83 topographic controls on drainage of surface and subsurface water (Devito et al., 2005). In low
84 elevation catchments across the North German Plain, streams are usually groundwater-
85 dominated, but the temporal and spatial heterogeneities in the hydrological functioning of these
86 catchments are still not fully understood (Boulton and Hancock, 2006). For example, there is still
87 a limited evidence base for quantifying how drought affects groundwater recharge and stream
88 flow generation in lowland areas in Central Europe, including the cessation of flow during the
89 summer (Germer et al., 2011).

90 To help address these knowledge gaps, here, we present a comprehensive set of ecohydrological
91 hydrometric and stable water isotope data of two years for the Demnitzer Mill Creek catchment,
92 Northeast Germany. The data set is unique in its integrative characteristics; that the different
93 compartments of the CZ were sampled across a mesoscale catchment in terms of their isotopic
94 signature and supporting ecohydrological data. By coincidence, these first two years, of what will
95 be a long-term study, captured the changing impacts of a prolonged drought period (2018-2020)
96 with a strong negative rainfall anomaly that became the most severe regional drought so far in
97 the 21st century (Kleine et al., 2021a). The data allow the effects of droughts (and their
98 persistence) on water storage, fluxes and age dynamics in the CZ to be investigated (Smith et al.,
99 2022). Our objective here is to provide this high spatio-temporal resolution ecohydrological
100 dataset to improve understanding of the storages and flow pathways of both blue and green water
101 across processes at the larger catchment scale in lowland catchments. We are continuing these
102 observations to assess long-term climatic trends at this drought sensitive region of Northeast
103 Germany, which is characterized by high water losses due to evapotranspiration and poor water
104 retention in the widespread sandy soils (Smith et al., 2021). Further, these data can potentially be
105 used to understand the hydrologic functioning of other drought sensitive regions beyond northeast
106 Germany.

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109 **2. Site description**
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 114 **Figure 1: The Demnitzer Mill Creek catchment and its location within Europe and**
 115 **Germany. Measurement types are indicated in the legend, with red indicating no isotope**
 116 **measurements, black and purple indicating isotope measurements, and purple**
 117 **additionally indicating sap flow and sap isotope measurements. Meteorological**
 118 **measurements at Neu Madlitz were conducted by the German Weather service (DWD**
 119 **Deutscher Wetterdienst).)**
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 121

122 **Table 1** - Overview of the properties of the Demnitzer Millcreek catchment at the catchment
 123 outlet. Overview includes physiological characteristics, landuse, and geology.

Area (km²)	66.39	Topographic Relief (m)	50.23
Runoff Ratio	0.10	Mean Slope (%)	1.98
Landuse (%)		Geology (%)	
Mixed Forest	1.0	Base moraine	35.5
Conifer Forest	29.2	End moraine	2.3
Broadleaf Forest	6.0	Deposits of glacial valleys	6.9
Peat	0.7	Peat Fen	5.9
Pasture	10.2	Periglacial/fluviol deposits	16.3
Agricultural/arable land	50.4	Glacial/fluviol deposits	31.1
Urban	2.5	Sandy peat fen	2.0

124
 125 The data presented here were monitored in the Demnitzer Millcreek catchment (DMC) located in
 126 NE Germany (52°23'N, 14°15'E; Figure 1). The DMC is a lowland drought-sensitive area south
 127 east of Berlin, the German capital, and situated in the North German Plain. The region has high
 128 socio-economic significance through the provision of numerous ecosystem services; including
 129 food security, timber production, groundwater recharge and river flow generation which sustains
 130 drinking water supplies for Berlin (Kleine et al., 2021a). The original motivation behind establishing
 131 DMC as an observatory in 1990 was to investigate the impact of agricultural pollutants on surface
 132 water quality (Gelbrecht et al., 2000, 2005).

133 The hydroclimate is temperate with warm, humid summers (Kottek et al., 2006). Mean annual
 134 precipitation and air temperature are 567 mm yr⁻¹ and 9.6°C, respectively (DWD, 2020, for 2006-
 135 2015). Seasonal contrasts are characterized by higher summer precipitation, mainly from high
 136 intensity, convective events; and slightly lower precipitation during frequent, frontal rainfall events
 137 in winter. The landscape was shaped by the last glaciation (Weichselian); soils are predominantly
 138 sandy and formed on glacial and fluvial deposits (Kleine et al., 2021b). The catchment is
 139 dominated by groundwater and likely had little surface runoff before human intervention.
 140 Previously, numerous peat fens and freshwater lakes in hollows existed, but these were drained
 141 during a long history of anthropogenic management (Nützmänn et al., 2011). Land use is currently
 142 dominated by farming and forestry (Kleine et al., 2020; Smith et al., 2020c). The catchment is also
 143 relatively sparsely populated, and has recently experienced recolonization of beaver (Smith et al.,
 144 2020a), wolf (Vogel, 2014) and even sporadic sighting of elk (Martin, 2014).

145 Maintenance of crucial ecosystem services in the landscape is dependent on sufficient seasonal
 146 precipitation input to sustain adequate soil moisture levels in the rooting zone to support crop and
 147 tree growth (Drastig et al., 2011); and acceptable groundwater recharge to sustain groundwater-
 148 surface water exchanges. However, high water losses due to evapotranspiration (~ 90 % of total
 149 precipitation), particularly from forested areas and poor water retention in the widespread sandy
 150 soils (Smith et al., 2021), result in catchment drought sensitivity (Kleine et al., 2020). Further,
 151 increased flow disconnections and fragmentation of the stream network occurs during droughts
 152 (Kleine et al., 2021a; Smith et al., 2021).

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154 **3. Data and instrumentation overview**

155 **3.1 Instrumentation overview**

156 A fully automatic weather station (AWS) was installed and has been operated in Hasenfelde (Hf,
157 Figure 1) since April 2018, including net radiation, air temperature, relative humidity, precipitation
158 and ground heat flux every 15-minutes. A modified autosampler (ISCO 3700, Teledyne Isco,
159 Lincoln, USA) was installed nearby to collect daily samples of precipitation to supplement the
160 AWS. Weekly cumulative precipitation was additionally collected at four locations nested from
161 north to south in the catchment: Marxdorfer St., Demnitz Mill, Bruchmill, and Berkenbruck (Figure
162 1&2) from July 2018 to April 2020. Throughfall was collected under the canopy at Forest A at five
163 locations (Forest A1-5) within a 10 m square fenced area. Throughfall was collected using
164 standard rain gauges (Rain gauge kit, S. Brannan & Sons, Cleator Moor, UK;
165 <https://doi.org/10.18728/igb-fred-623.0>)

166 Soil moisture and temperature profiles were established at Forest A (FA) and Grass A (GA) in
167 June 2018 with 18 sensors per site (SMT-100, Umwelt-Geräte-Technik GmbH, Müncheberg,
168 Germany). The sensors were distributed equally at soil depths of 20, 60, and 100cm at each site
169 (i.e. three sensors per depth), measuring every 15-minutes (<https://doi.org/10.18728/igb-fred-623.0>).
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171 Sap flow measurements were established in 12 trees at Forest A including Scots Pine (*Pinus*
172 *sylvestris*), European Oak (*Quercus robur*), common hazel (*Corylus avellana*), and Red Oak
173 (*Quercus rubra*). Measurements were conducted using 2-4 radially installed thermal dissipation-
174 based sap flow sensors (TDP probes, Dynamix Inc., Houston, TX, USA). Sap flow measurements
175 were recorded every 15 minutes (<https://doi.org/10.18728/igb-fred-623.0>).

176 Stream water level was established at four locations within the catchment; Peat North, Bruchmill,
177 Demnitz Mill, and Berkenbruck (<https://doi.org/10.18728/igb-fred-623.0>). Water level
178 measurements were established by IGB Leibniz Institute of Freshwater Ecology and Inland
179 Fisheries and recorded with divers (Micro 10m and Baro) at Peat North and Demnitz Mill, and at
180 Bruchmill (Van Essen Instruments). The divers utilized at each site include an internal
181 atmospheric pressure correction (AquiLite ATP 10, AquiTronic Umweltmeßtechnik GmbH,
182 Kirchheim/Teck, Germany). Water level measurements began at Demnitz Mill in 1986, and in
183 January and June 2018 for Peat North and Bruchmill, respectively. Water level has been recorded
184 since 1982 at Berkenbruck using pressure transducers and was established and collected by the
185 Landesamt für Umwelt. Channel stability at Demnitz Mill and Berkenbruck has permitted rating
186 curve development to translate water level measurements to discharge. Daily stream water
187 samples for stable water isotope analysis were also collected at Bruchmill from an autosampler
188 (ISCO 3700, Teledyne Isco, Lincoln, USA), which was established in December 2018
189 (<https://doi.org/10.18728/igb-fred-623.0>). Groundwater level divers were installed at five locations
190 throughout the catchment in 2001 (GW3, GW4, GW5, GW7, and GW8) (Figure 1&2).
191 Groundwater level at each site was measured every four hours with an AquiLite ATP-10 diver

192 (AquiTronic Umweltmeßtechnik GmbH, Kirchheim/Teck, Germany) with internal correction for
193 atmospheric pressure (<https://doi.org/10.18728/igb-fred-623.0>).

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195 **3.2 Isotope sampling overview**

196 Manual sampling from different locations and different water cycle / landscape compartments
197 supplemented the autosamplers installed for precipitation at Hasenfelde and for stream water at
198 Bruchmill. Samples were taken from the weekly cumulative precipitation and throughfall (Forest
199 A) for each location (Figure 2). Further, monthly samples of soil water were taken at 6 depths
200 (2.5, 7.5, 15, 30, 60, 90 cm) in triplicate for Forest A and Grass A. This was complemented by
201 synoptic, spatially distributed sampling of the upper 30cm in 2019. Samples were placed in a
202 sterile zip-lock bag (CB400-420siZ, Weber Packaging GmbH, Güglingen, Germany) and
203 analyzed using the direct water vapour equilibrium method (Wassenaar et al., 2008). Weekly grab
204 samples of stream water were taken at all nested stream water locations (eight locations; Fig 1.).
205 Groundwater isotopes were sampled at six groundwater wells (GW3, GW8, GW DA, GW6, GW
206 WLW, GW BB). Vegetation isotopic sampling was conducted by taking twig samples from different
207 vegetation in Forest A and samples of the non-green stem of the grass at site Grass A. Vegetation
208 samples were stored at -20°C after sampling until analysis. Reference for all isotope samples is
209 <https://doi.org/10.18728/igb-fred-623.0>.

210 A layer of paraffin was added to the bottom of all autosampler containers to prevent evaporation
211 and fractionation from collected water. Autosamplers are emptied each week. Collected weekly
212 precipitation, throughfall, stream water, and groundwater were sealed and refrigerated until
213 isotopic analysis (usually within one week).

214 All liquid water samples (isotopes in precipitation P_{iso} , in throughfall THR_{iso} , in streamwater Q_{iso} ,
215 in groundwater GW_{iso}) were filtered (0.2 μ m, cellulose acetate, Lab Logistics Group GmbH,
216 Meckenheim, Germany) and cooled before being analyzed using Cavity Ring-Down
217 Spectroscopy (CRDS, L2130-i, Picarro, Inc., CA, USA). Additionally, the CRDS was used for the
218 analysis of soil water extracted via the direct liquid water equilibrium method. Vegetation samples
219 were extracted in January 2020 using the cryogenic extraction method given in Dubbert et al.
220 (2013, 2014) and analyzed with the CRDS.

221

222 **Table 2 – Overview of site locations in DMC, including site name, coordinates, data**
 223 **collected, start and end dates, and resolution. N/A indicates not applicable, P is**
 224 **precipitation, GW is groundwater level, THR is throughfall, Ts is soil temperature, va is**
 225 **wind speed/direction, Ta is air temperature, Pa is air pressure, RH is relative humidity,**
 226 **NR is net radiation, Sap is sap flow, and subscript iso indicates isotopic sampling. AWS**
 227 **indicates measurements of P, va, Ta, Pa, RH, and NR**

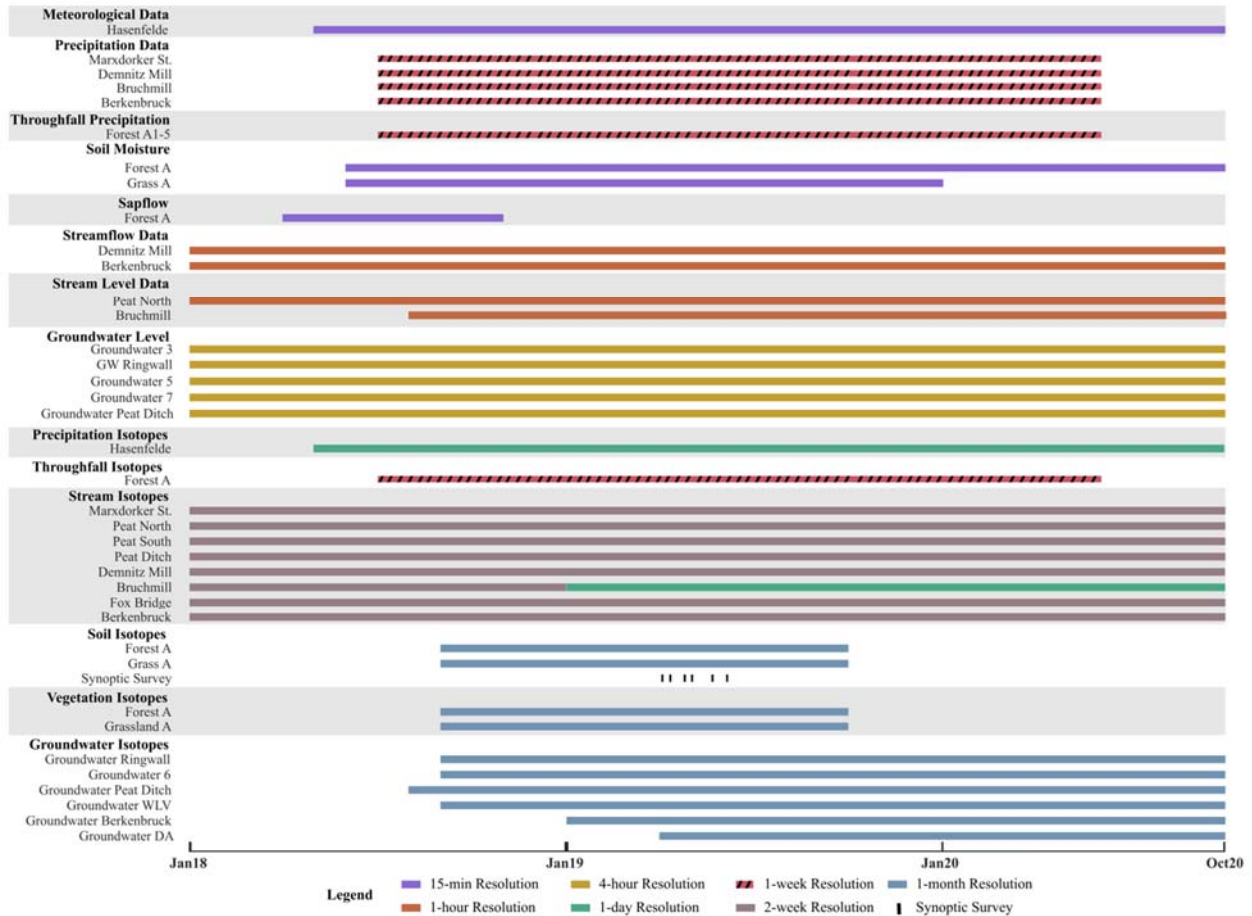
Site Name	ID	Location (UTM 33N)		Data Type	Installation/Start Date	Discontinued/End Date	Resolution	
		Latitude	Longitude				Temporal	Spatial
Marxdo rfer St.	Marxdo rfer St.	5810076	449773	P	Jul 9, 2018	Jun 2, 2020	Weekly	N/A
				P _{iso}				
				Q _{iso}	Jan 10, 2018	N/A		
				T _s	Aug 16, 2019	Jul 11, 2020	15-min	5cm
Hasenf elde	Hf	5809705	446068	P	Mar 17, 2018	N/A	15-min	2m
				P _{iso}	Jul 12, 2018		Daily	N/A
				va	Mar 17, 2018		15-min	2m
				Ta				
				Pa				
				RH				
				NR	Aug 16, 2019		Jul 11, 2020	5cm
T _s								
Ground water DA	GW DA	5808335	447527	GW _{iso}	Apr 16, 2019	N/A	Monthly	N/A
Peat North	PN	5807703	447474	Q _{iso}	Jan 10, 2018		Weekly	
Ground water 3	GW3	5807499	447582	GW	Jan 10, 2001		4-hour	
Ground water Ringwa ll	GW4	5807247	447233	GW	Feb 22, 2001		4-hour	
				GW _{iso}	Sep 11, 2018		Monthly	
Ground water 5	GW5	5807099	447490	GW	Jan 10, 2001		4-hour	
Peat Ditch	Peat Ditch	5806364	446487	Q _{iso}	Mar 21, 2018		Weekly	
Ground water Peat Ditch	GW8	5806320	446488	GW	Jan 10, 2001		4-hour	
				GW _{iso}	Aug 15, 2018		Monthly	
Ground water 7	GW7	5806307	447726	GW	Feb 22, 2001		4-hour	
Ground water 6	GW6	5806274	447678	GW _{iso}	Sep 11, 2018	Monthly		
Peat South	Peat South	5806262	447712	Q _{iso}	Jan 10, 2018	N/A	Weekly	N/A
				T _s	Aug 16, 2019	Jul 11, 2020	15-min	5cm
Forest A	FA	5805520	445731	Sap	Apr 21, 2018	Nov 1, 2018	15-min	12 Trees
				SM	Jun 15, 2018	N/A		6 sites, 20, 60, 100cm depths

				SM _{iso}	Oct 18, 2018	Jul 16, 2019	Monthly	N/A
				THR	Jul 11, 2018	May 19, 2020	Weekly	5 sites
				THR _{iso}				
				T _s	Jun 15, 2018	N/A	15-min	6 sites, 20, 60, 100cm depths
Grass A	GA	5805125	445495	SM	Jun 15, 2018	Jan 7, 2020	15-min	6 sites, 20, 60, 100cm depths
				SM _{iso}	Oct 18, 2018	Jul 16, 2019	Monthly	N/A
				T _s	Jun 15, 2018	Jan 7, 2020	15-min	6 sites, 20, 60, 100cm depths
Bruchmill	Bruchmill	5805088	445459	P	Jul 9, 2018	Jun 2, 2020	Weekly	N/A
				P _{iso}				
				Q _{iso}	Jan 10, 2018 (weekly) Dec 28, 2018 (daily)	Dec 28, 2018 (weekly)	Weekly / Daily	
Ground water WLW	GW WLW	5803322	445982	GW _{iso}	Sep 20, 2018	N/A	Monthly	
Demnitz Mill	Demnitz Mill	5802298	445188	P	Jul 9, 2018	Jun 2, 2020	Weekly	
				P _{iso}				
				Q	Feb 22, 2011	N/A	4-hour	
Q _{iso}	Jan 10, 2018	Weekly						
Fox Bridge	Fox Bridge	5801469	444189	Q _{iso}	Jan 10, 2018	N/A	Weekly	
Ground water Berkenbruck	GW BB	5799862	444611	GW _{iso}	Jan 21, 2019		Monthly	
Berkenbruck	Berkenbruck	5799604	444737	P	Jul 9, 2018	Jun 2, 2020	Weekly	N/A
				P _{iso}				
				Q	Nov 1, 1982	N/A	Daily	
				Q _{iso}	Jan 10, 2018		Weekly	
T _s	Aug 16, 2019	Jul 11, 2020	15-min	5cm				

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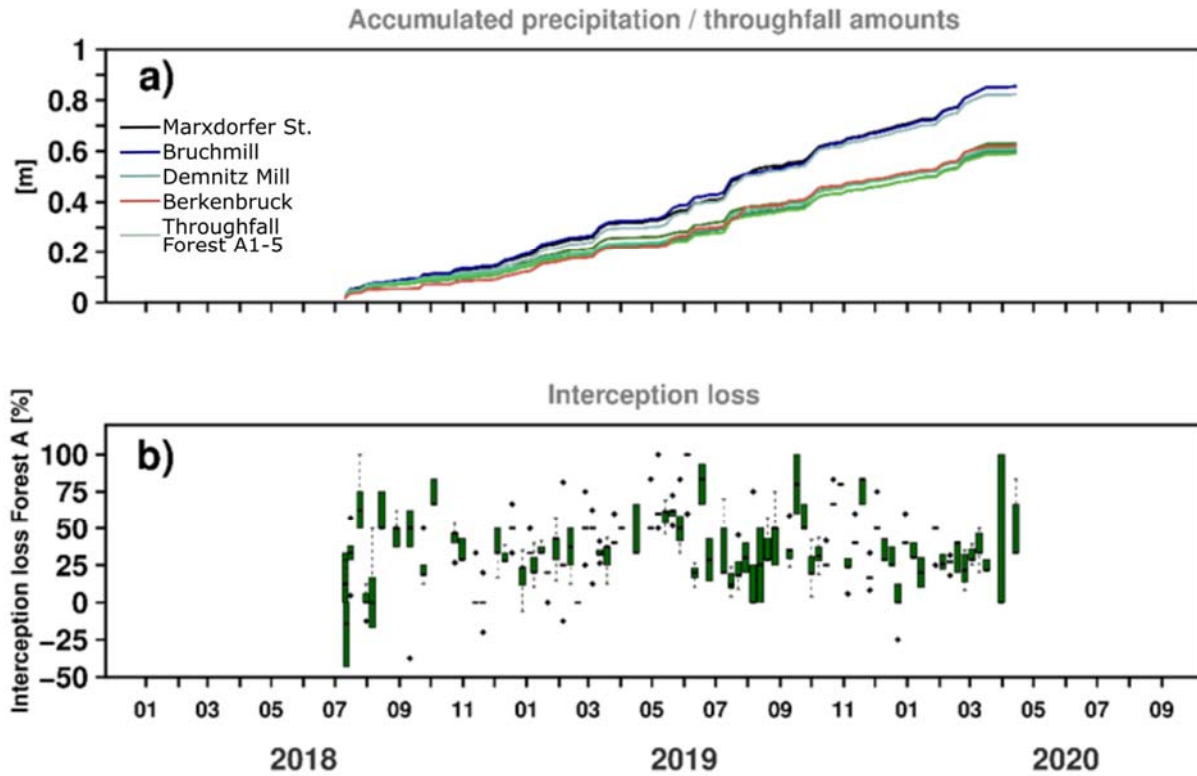


231
 232 **Figure 2: Measurement period for each parameter at each site and temporal resolution**
 233 **(colour code) of the measurements within the Demnitzer Millcreek Catchment including**
 234 **meteorological, soil, vegetation, stream, and groundwater hydrological and isotope data**
 235 **sets.**

236 **4. Precipitation and throughfall data**

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238 Monitoring for precipitation commenced in the 2018 summer drought when low rainfall inputs
239 continued through the following winter (Figure 3a). Large rainfall events (>20 mm/d) were
240 relatively rare and mostly summer convective storms. Even by summer 2020, most months had
241 below average rainfall. Throughfall at the Forest A site typically was 70-90 % of incident rainfall,
242 with higher interception losses in low intensity summer storms and lowest in winter or high
243 intensity summer storms. Heterogeneity in throughfall was marked (Figure 3b), emphasizing the
244 importance of the forest canopy in redistributing net rainfall to the forest floor.

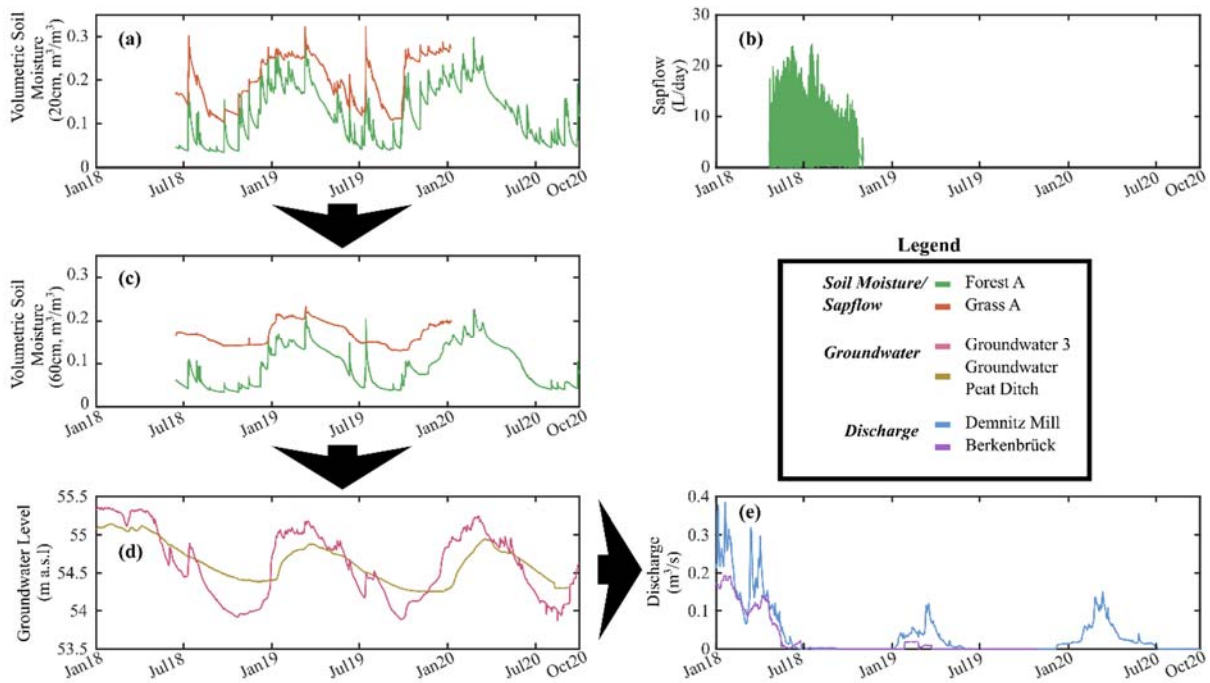


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246 **Figure 3: (a) Cumulative precipitation and throughfall at multiple locations throughout**
247 **the catchment. Throughfall was collected weekly at Forest A with (b) five samplers (1-5)**
248 **distributed throughout the 10m square fenced region.**

249 **5. Catchment hydrologica data**

250 Rainfall fluxes mostly drove short term soil moisture variations (Figure 4a, c); which were more
251 responsive in the upper soil layers (at 20 cm) than deeper layers. There was higher variability in
252 volumetric soil moisture under forested land cover, where soils are sandier, more structured and
253 effective rainfall is lower due to interception losses. Seasonality in evapotranspiration (usefully
254 indexed by sapflow in Figure 4b) modulated the effects of rainfall on soil moisture storage.
255 Seasonal soil moisture dynamics also governed groundwater recharge and variation in
256 groundwater levels, which had an annual range of ~1.5 m at well G3 and ~1m at the peat ditch
257 well (Figure 4d). Despite clear winter recharge and spring drawdown in each well, peak winter
258 and summer levels were lower in 2019 and still in 2020 despite a slight recovery compared to
259 2018 indicating the cumulative “memory effects” of the drought. This was also evident in the
260 stream hydrograph with very low discharge peaks in 2019 and 2020, which also had prolonged
261 periods where flow ceased in the summer, particularly at Berkenbrück. Thus, winter soil moisture
262 replenishment was insufficient to match long-term groundwater recharge. These different
263 correlations underline the added value of simultaneous data from long-term study sites on
264 transpiration, soil water, groundwater and stream flow as droughts develop (Smith et al., 2022).
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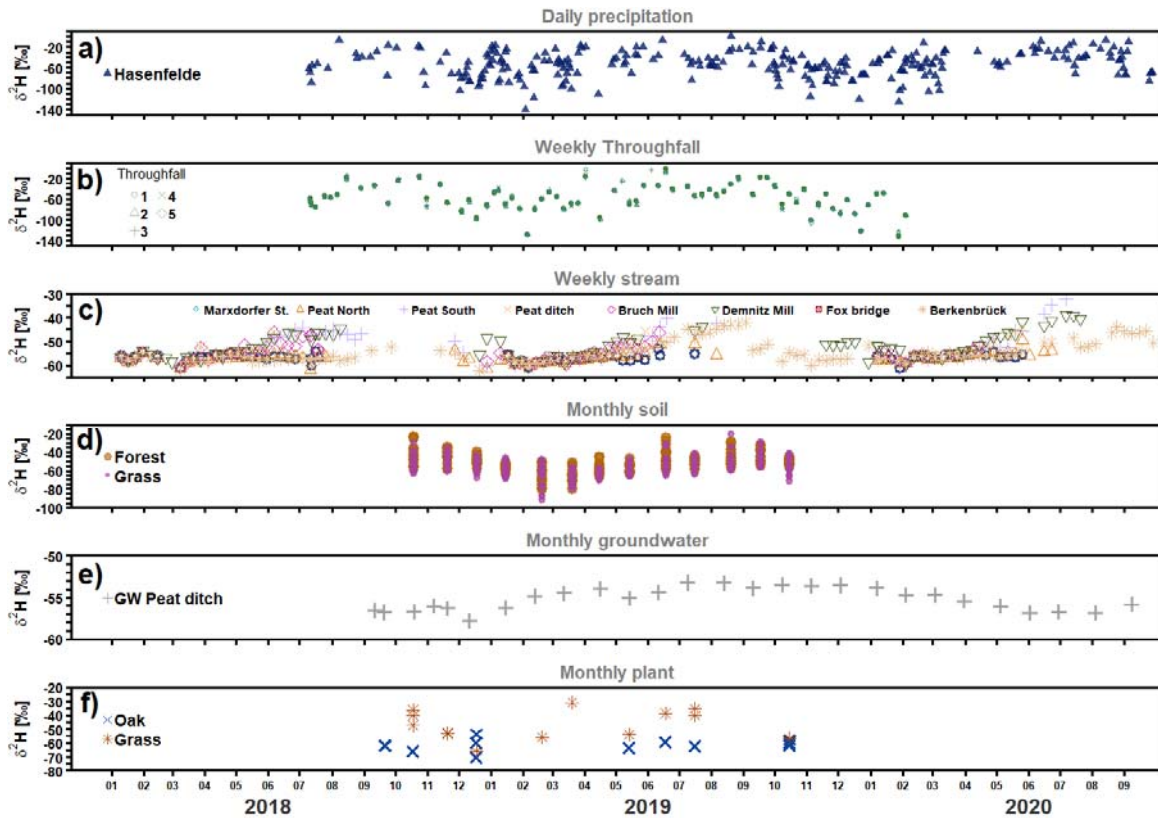


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268 **Figure 4: (a) Shallow soil moisture, (b) sapflow, (c) deep soil moisture, (d) groundwater**
269 **levels and (e) discharge within the Demnitzer Millcreek catchment. Arrows show**
270 **connections between layers and fluxes. *Groundwater 3 is within the wetland and**
271 **Groundwater Peat Ditch is outside the wetland (near Forest A and Grass A, Fig. 1).**

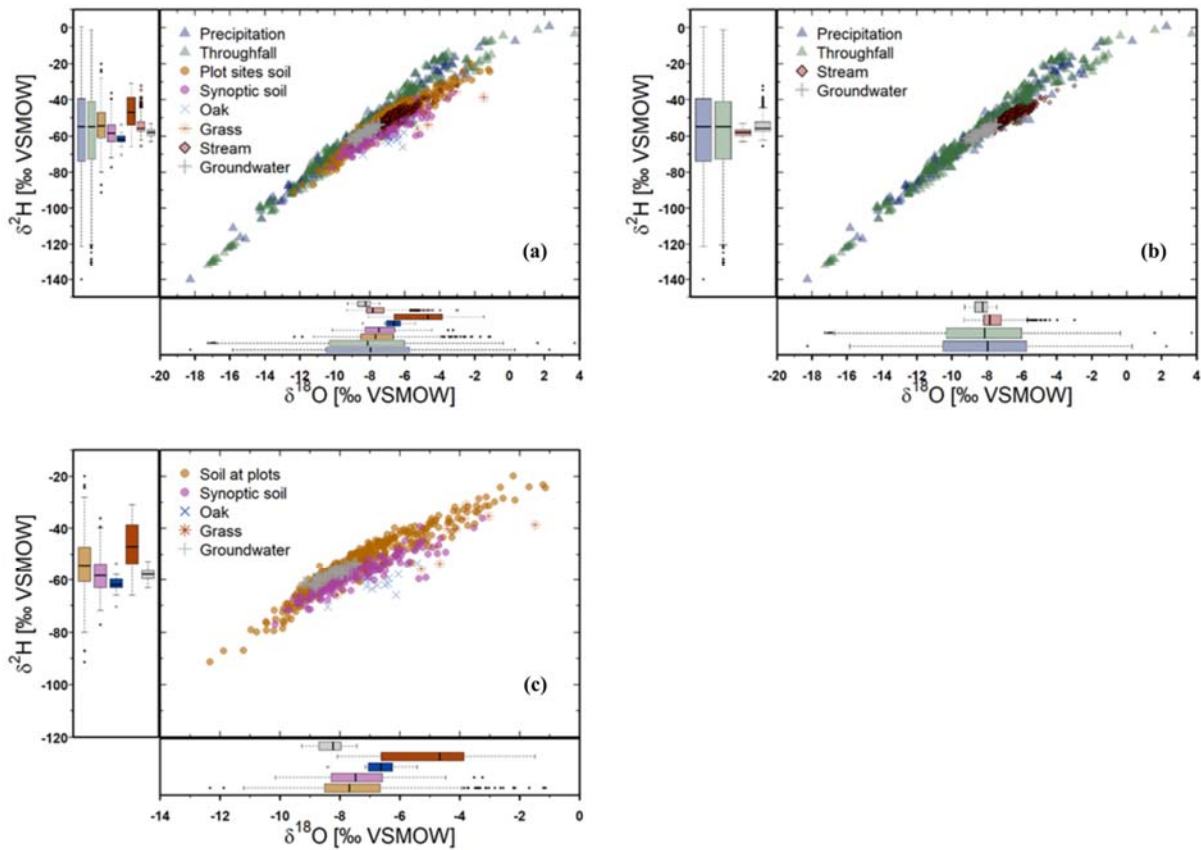
272 **6. Stable water isotopes**

273 Stable water isotope signatures in precipitation showed high day-to-day variability superimposed
274 on strong seasonality; with more depleted values in winter and more enriched values in summer
275 (Fig 5a). Interestingly, weekly throughfall signatures were very similar to the (weekly and daily)
276 precipitation signal showing no strong signs of evaporative fractionation during canopy storage
277 (Fig 5b). This likely reflects the high intensity nature of most summer rainfall, which affords limited
278 opportunity for canopy evaporation. Streamwater signatures at all nested sites showed similar
279 seasonality but much more damping in the signal (Fig. 5c). Groundwater was most damped, and
280 similar in composition to streamflow during winter (Fig 5d). In summer, sites downstream of
281 Marxdorfer Strasse showed evidence of evaporative fractionation from either the channel network
282 or riparian soils and plotted below the meteoric water line before stream flow ceased. Monthly soil
283 water samples showed higher variability in isotopic composition under forest than under grass,
284 mainly reflecting soil characteristics with more retentive, loamy and wetter soils at the grassland
285 site buffering the effects of rainfall inputs. At both sites, seasonal variation in isotopic composition
286 tracked precipitation, though in deeper soil, the isotopic signal was more damped. Vegetation
287 samples from the oaks showed higher variation than from grass.
288



289 **Figure 5: Time series of deuterium ($\delta^2\text{H}$) in (a) precipitation, (b) throughfall (Forest A), (c)**
290 **stream water, (d) soil water, (e) groundwater and (f) plant samples at various locations in**
291 **the catchment.**
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Figure 6: Dual isotope space ($\delta^{2}\text{H}$ - $\delta^{18}\text{O}$) plots for (a) all measured isotopic datasets, (b) precipitation, throughfall, stream, and groundwater, and (c) soil (multiple depths), synoptic soil survey (upper 30cm), vegetation, and groundwater.

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Differences in the isotope dynamics of different critical zone compartments are shown in dual isotope space in Figure 6a. The damping of precipitation in groundwater and streamflow is apparent, as is the fractionation of more enriched summer stream flow samples (Figure 6b). The role of the soil in partitioning water is apparent from the overlap between deeper soil horizons and groundwater which were both more weighted to winter precipitation – when recharge is greatest (Fig 6c). Xylem water in oaks and grass tended to show the effects of fractionation, which was most marked in the oaks and may point to different soil water sources of root uptake.

307 **7. Data availability**

308 All data presented in this paper are available from the IGB open data repository FRED
309 <https://fred.igb-berlin.de/data/package/622> (Tetzlaff et al., 2022). The data are published with
310 detailed metadata (<https://doi.org/10.18728/igb-fred-623.0>) and contact information for any further
311 questions. There is a readme section per each dataset. We also included a digital elevation model,
312 shapefile of the catchment boundary and the station locations.

313

314 **8. Summary**

315 The integrated data set presented in this paper is unique because: (1) it captures complicated
316 ecohydrological dynamics over two years during an exceptional drought (in 2018/2019) in Central
317 Europe; (2) the different compartments of the critical zone were monitored through stable water
318 isotope data and complimentary ecohydrological data for contrasting land use and (3) multi-scale,
319 nested catchment time series were derived. In total data from 49 time series / data sets are
320 available. The data are quality controlled. We included meteorological data and precipitation and
321 throughfall amount. Catchment response data include stream discharge at the catchment outlet
322 and another nested site, and stream level data at two further sites; soil moisture from multiple
323 depths at two locations (two different landuses), groundwater level data at five locations and
324 sapflow measurements from one forest location. Stable water isotope data include precipitation
325 water, throughfall, streamwater at eight sites, soil water isotopes from two sites plus spatially
326 distributed samples of upper soils, vegetation samples at two locations and groundwater at six
327 locations. Data continue to be collected and updated data sets will be published based on
328 available resources.

329 As such, these data provide an excellent, integrated ecohydrological perspective on the drought
330 response of a lowland agricultural landscape. Such data are of course important in their own right,
331 but are equally invaluable for challenging environmental models as constraints on internal model
332 function that can be used to increase confidence in the use of models in projecting the impacts of
333 future change. Integrated data like the ones summarised here are also important for a range of
334 scientific questions that are growing in importance as the effects of climate change become more
335 apparent. These include understanding how do droughts develop and propagate through
336 components of hydrological systems and compartments of the critical zone? What are the effects
337 of land cover on this propagation and how does it affect water cycling in vegetation? How long
338 does recovery of different system components take once rainfall anomalies become positive?
339 How resilient are different critical zone compartments or entire landscapes against climate
340 extremes such as droughts? Hopefully, this data set will be used by scientists to increase
341 understanding on critical issues such as what are the water footprints of alternative land uses and
342 how can these be reduced whilst maintaining societal needs. This will help to contribute to the
343 development of more sustainable and resilient land and water management policies that will be
344 needed in the face of increased longevity and frequency of droughts.

345

346 **Author contributions:** AS and LK prepared the data sets. Datasets were collected by LK and
347 JF. Isotope data were analysed by DD. DT, CS, AS prepared the manuscript with contributions
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349

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351

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354

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360

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