

Enhanced automated meteorological observations at the Canadian Arctic weather science (CAWS) supersites

Zen Mariani¹, Laura Huang¹, Robert Crawford¹, Jean-Pierre Blanchet², Shannon Hicks-Jalali¹, Eva Mekis³, Ludovick Pelletier², Peter Rodriguez¹, and Kevin Strawbridge⁴

5

¹Meteorological Research Division, Environment and Climate Change Canada, Toronto, M3H 5T6, Canada

²Centre ESCER, Département des sciences de la Terre et de l'atmosphère, Université du Québec à Montréal, H2L 2C4, Canada

³Climate Research Division, Environment and Climate Change Canada, Toronto, M3H 5T6, Canada

⁴Air Quality Research Division, Environment and Climate Change Canada, Toronto, M3H 5T6, Canada

10 *Correspondence to:* Zen Mariani (zen.mariani@ec.gc.ca)

Abstract. The changing Arctic climate is creating increased economic, transportation, and recreational activities requiring reliable and relevant weather information. However, the Canadian Arctic is sparsely observed and processes governing weather systems in the Arctic are not well understood. There is a recognized lack of meteorological data to characterize the Arctic atmosphere for operational forecasting and to support process studies, satellite calibration/validation, search and rescue operations (which are increasing in the region), high impact weather (HIW) detection and prediction, and numerical weather prediction (NWP) model verification and evaluation. To address this need, Environment and Climate Change Canada commissioned two supersites; one in Iqaluit (63.74°N, 68.51°W) in September 2015 and the other in Whitehorse (60.71°N, 135.07°W) in November 2017 as part of the Canadian Arctic Weather Science (CAWS) project. The primary goals of CAWS are to provide enhanced meteorological observations in the Canadian Arctic for HIW nowcasting (short-range forecast) and NWP model verification, evaluation, and process studies, and to provide recommendations on the optimal cost-effective observing system for the Canadian Arctic. Both sites are in Provincial/Territorial capitals and are economic hubs for the region; they also act as transportation gateways to the North and are in the path of several common Arctic storm tracks. The supersites are located at or next to major airports and existing Meteorological Service of Canada ground-based weather stations that provide standard meteorological surface observations and upper air radiosonde observations; they are also uniquely situated in close proximity to frequent overpasses by polar-orbiting satellites. The suite of in-situ and remote sensing instruments at each site are completely automated (no on-site operator) and operate continuously in all weather conditions, providing near-real time data to operational weather forecasters, the public, and researchers via obsr.ca. The two sites have similar instruments, including mobile Doppler weather radars, multiple vertically-profiling and/or scanning lidars (Doppler, ceilometer, water vapour), optical disdrometers, precipitation gauges in different shielded configurations, present weather sensors, fog monitoring devices, radiation flux sensors, and other meteorological instruments. Details on the two supersites, the suites of instruments deployed, data collection methods, and example case studies of HIW events are discussed. CAWS data are publically accessible via the Canadian Government Open Data Portal (<https://doi.org/10.18164/ff771396-b22c-4bc3-844d->

38fc697049e9 (Mariani et al., 2022a) and <https://doi.org/10.18164/d92ed3cf-4ba0-4473-beec-357ec45b0e78> (Mariani et al., 2022b)); this dataset is being used to improve our understanding of synoptic and fine-scale meteorological processes in the Arctic and sub-Arctic, including HIW detection and prediction and NWP verification, assimilation, and processes.

1 Introduction

Economic activity in the Arctic is growing due to increasing population, transportation, tourism, and resource development with the opening of the North-west passage. For instance, marine and air traffic have significantly increased in the region (Smith and Stephenson, 2013; Arctic Council, 2017). At the same time, the changing climate, which is amplified in the Arctic region, induces changes in weather events with a high socio-economic impact (WMO, 2011). A disproportionate number of Search and Rescue (SAR) incidents occur in Canada's northern territories relative to the southern provinces (Government of Canada, 2016; Statistics Canada, 2016). Increasing demands will be placed on transportation and SAR-related infrastructure and services as high-impact weather (HIW) conditions are expected to become more frequent, longer in duration, and less predictable in the future (Ford et al., 2013). The provision of meteorological observations can help individuals, groups and organizations make informed decisions about when to safely travel, conduct particular activities, and take precautionary or protective actions. Such observations can reduce weather vulnerability, improve HIW warnings, prevent SAR incidents from occurring, and support SAR operations when undertaken (WMO, 2017).

Given the sparse availability of meteorological data in the Arctic, operational weather forecasters rely heavily on output from numerical weather prediction (NWP) models. Unfortunately, almost all international NWP models exhibit poor performance >60° N with significant errors in forecasted pressure and winds (Cassano et al., 2011; Schyberg and Randriamampianina, 2015; Riishojgaard, 2015). The primary cause of these errors is the large geographic gaps in meteorological measurements; despite Canada encompassing roughly 40% of the entire Arctic region with > 200,000 inhabitants (global), there exist only seven upper air stations (profile observations above the surface) and no weather radar data. The Canadian Network for the Detection of Climate Change research site at Eureka, NU, (80.05°N, 86.42°W) is equipped with remote sensing meteorological and climate observations (e.g., Lesins et al., 2009); otherwise the few ground-based weather stations that exist in the Arctic only provide standard surface meteorological observations (surface pressure, temperature, humidity, and wind). An overview of the few previous/current Arctic science projects that provide enhanced meteorological or climatological observations is provided in Joe et al. (2020).

As NWP model resolution increases, there exists a growing need for high spatial- and temporal-resolution meteorological measurements in the Arctic beyond the standard surface measurements. Such observations can be used to validate, inter-compare, and perform NWP process studies which can eventually lead to changes that enhance the performance of NWP systems; this is one of the foci of the World Meteorological Organization's (WMO) Year of Polar Prediction (YOPP) project

65 (core phase: mid-2017 to mid-2019) (Koltzow et al., 2019). Validation of NWP output within the planetary boundary layer (PBL) is particularly essential since the representation of the PBL's structure and physical processes in the Arctic remain a challenge in NWP systems (Cassano et al., 2011; Illingworth et al., 2015; Schyberg and Randriamampianina, 2015).

To address this need, Environment and Climate Change Canada (ECCC) commissioned two supersites at Iqaluit (airport designator: CYFB, 63.74°N, 68.51°W, 11 m a.s.l.) and Whitehorse (airport designator: CYXY, 60.71°N, 135.07°W, 682 m a.s.l.). **The two sites are representative of their regions and provide contrasting conditions: e.g., Western vs. Eastern Arctic, mountainous vs. tundra, and inland valley vs. marine.** These sites are representative of their corresponding regions.. Both sites were designated as official YOPP supersites during the entire YOPP project (including pre- and post-YOPP phases). The sites provide fully automated and continuous observations of vertically-resolved winds, water vapour, clouds and aerosols, as well as surface/soil observations of visibility, radiation fluxes, and precipitation during all weather conditions as part of the Canadian Arctic Weather Science (CAWS) project **(Joe et al., 2020). The new profiling observations of winds and water vapour, for instance, are crucial to determine fluxes of water vapour transport, the presence of atmospheric rivers, and hazardous wind conditions for aviation. Such profiling observations do not currently exist in the Arctic (except for standard radiosondes every 12 hr); as such these profile observations provide novel data useful for satellite calibration/validation, evaluating and improving NWP model performance above the surface layer, HIW classification (e.g., depth and height of blowing snow during a blizzard), short-term weather forecasting (nowcast), and for cloud microphysics studies.** The sites also conducted standard WMO surface meteorological observations at pre-existing co-located Meteorological Service of Canada (MSC) weather stations.

85 In the design of the CAWS supersites, emphasis was placed on deploying new remote sensing technologies that were fully automated to reduce operational costs and eliminate requirements for on-site personnel. The advantage of relying on remote sensing instruments to fill data gaps (in both time and vertically in space) for operational forecasters and improve NWP models is outlined in Illingworth et al. (2015). These new technologies underwent thorough multi-year evaluations in order to advise on a cost-effective Arctic observing system, a primary goal of the CAWS project (e.g., Mariani et al., 2020a; Mariani et al., 2020b; Mariani et al., 2021); these are the first multi-year evaluations to occur in the Arctic region for many of these instruments, whereas previous evaluations occurred over shorter periods at mid-latitudes (e.g., Kumer et al., 2014; Paschke et al., 2015; Newsom et al., 2020).

95 A wide range of meteorological conditions were observed at unprecedented resolution at the CAWS supersites. HIW events were frequently observed, including prolonged stratified wind and water vapour layers (Mariani et al., 2018), blizzards and low-visibility conditions (22% of days in Iqaluit experienced visibility < ½ standard mile during the study period), and more; these events impacted local communities in different ways, including the closure of airports, SAR efforts, and preventing

manual in-situ observations from taking place (e.g., radiosondes could not be launched at Iqaluit 13% of the time due to high surface winds).

100

The data collected during the CAWS project serves as ECCCC's primary contribution to providing enhanced meteorological observations during YOPP. Several other supersites from other meteorological agencies also contributed to YOPP; combined, the data collected at these supersites provides the most detailed pan-Arctic observational dataset for NWP evaluation to date. CAWS observations were/are provided to operational forecasters for nowcasting (short-range forecast) purposes, researchers, and the public in near-real time via the website obrs.ca. The data is also used to support informed decisions on NWP forecast model development and weather forecasting programs, and to enable ground-based calibration and validation of meteorological satellites, such as the ADM-Aeolus (e.g., Chou et al., 2021), GPM, and the upcoming EarthCARE and AOS satellite missions.

105

110

This paper describes the suite of instrumentation deployed to the two supersites, some of which are new state-of-the-art pre-production commercial units. The data collected at the two supersites fill crucial gaps in Arctic observations, particularly for upper-air (PBL) observations. Section 2 describes the two study areas and their climatology. Section 3 outlines the instrumentation used and datasets collected. Section 4 provides examples of observations at the two supersites in two case studies. Sections 5 and 6 provide details of the online database and concluding remarks, respectively.

115

2 Supersite Descriptions

2.1 Iqaluit

120

Iqaluit is the capital of the Territory of Nunavut with a population of over 8,000 inhabitants. It is the primary gateway for air and sea traffic for the central and Eastern Arctic; it is near many current and planned primary transportation corridors for marine vessels. As such, Iqaluit is commonly referred to as the 'gateway to the North.' The CAWS supersite is located ~200 m from the airport runway on existing MSC weather station property (Figure 1). All instruments are co-located to within (maximum) 140 m of each other on flat **permafrost terrain (rock / soil).**

125

Iqaluit is influenced by a diversity of synoptic storms originating from across the Arctic. Most typical storm tracks originate over the Western Arctic or the Prairies. These storms can produce very strong Easterly winds within the PBL that, despite Iqaluit's dry climate (< 200 mm annual precipitation), can cause blowing snow that severely reduces visibility during non-summer months. During the summer, the frequent formation of fog around the Frobisher Bay area also acts to limit visibility. Iqaluit experiences a wide range in surface temperatures (typically -35 to 20 °C) year-round with almost 21 hours of sunlight/darkness during Polar Day/Night. The city itself is located along the coast in a valley that runs in the NW to SE

130 direction; thus the primary direction of surface winds follows this direction. The surrounding region is relatively flat Arctic
tundra except for nearby hills (~300 m a.s.l.) approximately two kilometers to the NE of the supersite.

2.2 Whitehorse

135 Whitehorse is the provincial capital of the Yukon Territories with a population of over 26,000 inhabitants. Similar to Iqaluit,
it is the primary gateway for air traffic for all of the Yukon Territories, parts of Alaska, and the Western Arctic. The CAWS
supersite is located on the Erik Nielsen Whitehorse International Airport property, which itself is situated on a plateau
overlooking (~50 m above) the city (Figure 1). The supersite's instruments are installed on an elevated **wooden platform, all
within a few metres of each other above compact gravel**, while the MSC weather station is located off-site, 2.9 km NNW of
the airport.

140 Most storm tracks that pass through Whitehorse originate from the Eastern Pacific or over Alaska. The complex mountainous
terrain in this region strongly influences these systems; for instance, blocking systems from entering the valley and causing
leeside (upslope) precipitation. Contrary to Iqaluit, Whitehorse is located in a wide valley of the Yukon River with the Yukon
Ranges to its West (~1.6 km a.s.l. mountain peak) and East (~1.4 km a.s.l. mountain peak). Similar to Iqaluit, the primary
145 surface wind direction follows the valley (NNW). Despite its sub-Arctic location, it has a relatively dry climate with annual
precipitation < 270 mm and experiences an even wider range of temperatures (typically -30 to +30 °C) year-round.

3 Data Collection

3.1 Iqaluit Supersite Instrumentation

150 A suite of ground-based remote sensing and in-situ instruments were deployed to the Iqaluit supersite over a three-year period
starting in September 2015 as part of the CAWS project. All instruments were fully automated, operated 24/7 without needing
an operator at the site, and conducted new meteorological observations of variables not observed by the pre-existing MSC
standard surface meteorological observations. Data collection was continuous except for brief power outages that occurred
infrequently (about once a month) at the site or instrument-specific failures that required a remote reboot. **All instruments have
155 technical performance ratings suitable for Arctic conditions; as such, they are equipped with heaters, fans, and wiper blades to
remove accruing ice/snow. Remote monitoring via 4K cameras also enabled visual confirmation of the absence of snow or ice
accumulation (image archive is available upon request).** A complete list of all instruments and their locations is provided in
Table 1; images of the site and its instruments are provided in Figures 2 and 3.

160 Unless stated otherwise, all data files are in standard ASCII text file output formats developed by the instrument manufacturer and are easily readable. Data collection for most instruments is ongoing, though in limited capacity due to issues with travel related to COVID-19. An example of some of the enhanced meteorological observations collected between 2015 and 2021 by the CAWS instruments is provided in Figure 4; these observations illustrate the large range of different weather conditions observed during the majority of the entire study period.

165

A second, smaller site named “T121” is located on NavCanada property on a ridge on the city’s perimeter. The instruments at T121 overlook the airport and rest of the city (Figure 2). It sits atop a 170 m high ridge 2.28 km NNE of the main supersite. Since T121 is located above and outside of the valley, the meteorological data collected there provide context for the synoptic conditions surrounding the city.

170

3.1.1 Meteorological Service of Canada Weather Station

The MSC weather station has been in operation at Iqaluit since 1953. The building is managed by the Observing Systems and Engineering Division in MSC and acts as the central hub, connecting all instruments and related infrastructure at the Iqaluit supersite. The weather station conducts WMO-standardized hourly surface meteorological observations of surface temperature, relative humidity, pressure, wind speed and direction, and precipitation (Joe et al., 2020). This surface data is also available in near-real time at weather.gc.ca. Vaisala RS92 (Vaisala, 2007) and, after 2018, GRAW DFM-09 (GRAW, 2020) radiosondes were launched by an MSC operator twice a day (00:00 and 12:00 Coordinated Universal Time (UTC)) as per WMO guidelines at the Iqaluit weather station (WMO station code 71909). Radiosondes provide highly-accurate vertical profile observations of atmospheric temperature, relative humidity, pressure, wind speed and direction, and other parameters up to ~40 km a.g.l.. Meteorological Reports (METAR) were reported at the Iqaluit airport every hour and occasionally in between hours when conditions warranted a special (manual) report; they provide additional meteorological information including weather type, cloud amount, and cloud height in 3 layers.

180

3.1.2 Surface Visibility and Precipitation Type

185 Two Vaisala PWD52 Present Weather Detectors were deployed; one at the Iqaluit supersite and the other at the T121 site. The instruments operated continuously with limited data gaps. They provide observations of visibility, precipitation rate and type, and luminescence. Data was output in Vaisala’s Data Message 7 format (see user manual). Forward-scatter present weather detectors are ideally suited for observing Arctic conditions as they have greater sensitivity and respond to light precipitation conditions better than unshielded weighing gauges. Their sensitivities and applications are also well characterized (Barthazy and Schefold, 2006; Battaglia et al., 2010; Liu et al., 2013; Tokay et al., 2014; Zhang et al., 2015; Gultepe et al., 2016, 2017).

190

While these instruments have enabled more accurate estimates of light precipitation (see Table 1, ‘accuracy’), they are inherently limited to point measurements.

195 The PWD52 (and FS11P used in Whitehorse) meet Federal Aviation Administration and International Civil Aviation Organization specifications. Precipitation type and intensity are estimated based on an optical principle via the attenuation of a laser beam by falling particles. The precipitation type can be estimated by using empirical relationships between the observed diameter and fall speed of the particles (Gunn and Kinzer, 1949). Default settings for the precipitation intensity limits define the light (< 2 mm/hr), moderate (2-8 mm/hr), and heavy (> 8 mm/hr) precipitation flags reported in the data (different thresholds are used for snow). The precipitation classification algorithm is proprietary to the manufacturer (Vaisala) and was
200 used without modification.

3.1.3 Particle Imaging Package

The Particle Imaging Package (PIP) is a video disdrometer designed and built by NASA. It consists of a high-speed video camera (380 frames/second) with a 640 x 480 pixel charge-coupled device image sensor. It operated continuously with limited
205 data gaps. This camera is aimed at a bright (150 W) halogen lamp two metres away, resulting in an image resolution of ~ 0.1 mm x 0.1 mm. The PIP setup is unique in that precipitation particles are unimpeded by the instrument itself. Hydrometeor shadows are recorded as they fall through the observation volume; this enables observations of particle imagery used to calculate particle size distributions (PSDs), fall speed estimates, droplet size distributions (DSDs), precipitation rate, and density estimates continuously with 1-min resolution (Newman et al. 2009; Tiira et al. 2016; von Lerber et al. 2018; Pettersen
210 et al., 2020; Pettersen et al., 2021).

3.1.4 Radiation Flux Sensor Suite

Short and longwave radiation flux sensors were deployed to the Iqaluit supersite to characterize the radiative budget at the site. It was installed in September 2018 and operated continuously from that point onwards with limited data gaps. The flux sensor
215 suite consists of a 4 m mast with a sensor cross-arm attached near the top. The ends of the cross-arm point in the four cardinal directions (N, E, S, W). Two Kipp and Zonen CMP10 pyranometers (facing up and down) and six CGR4 pyrgeometers (facing up, down, and horizontally in the N, E, S, W directions) were installed on the mast’s crossarms to provide measurements of short and longwave radiation, respectively. The horizontal longwave sensors are a unique feature of the flux sensor suite; they are used to investigate horizontal longwave radiation fluxes at the site due to surrounding buildings and topography. All sensors
220 were equipped with CVF4 ventilation units to prevent fog/frost forming on the sensor’s dome, improving data quality and reliability. All sensors were tested in an environmental chamber for extreme cold in March 2018 at the ECCC Downsview Lab before being deployed to Iqaluit to ensure the units could perform nominally during the harsh climate in the Canadian Arctic.

225 Despite this, the observations should be treated with caution since the absence of ice/frost/snow on the radiometer's dome was not verified for every observation as it was in Cox et al. [2021]. A Campbell Scientific CR1000X data logger and CMD-A108 8-channel analog input module were used to record data.

230 In addition to the radiation fluxes, two Campbell Scientific SR50ATH snow depth sensors and a CS655 soil water content reflectometer with a soil temperature sensor were also installed. They provide observations of snow depth, soil moisture, and soil temperature below the flux sensor suite to further help characterize the site's radiative budget. Two flat calibration target pads were installed under each SR50ATH to ensure snow depth measurements were calibrated and recorded on a standardized surface. Finally, a Rosemount icing detector provides an indication of icing conditions (i.e., the presence of super-cooled water and an estimate of its quantity). It consists of a piezoelectric sensor that detects changes in its natural vibration frequency due to ice build up. As such, it is useful for determining whether ice and/or frost formed on/near the surface.

235 3.1.5 Far Infrared Radiometer

The Far Infrared Radiometer (FIRR) measures the downwelling long-wave far infrared radiation emitted by the atmosphere using newly developed microbolometer technology. The Iqaluit FIRR is a second-generation infrared radiometer developed by LR Tech Inc. based on its earlier version (Libois et al., 2016; Libois and Blanchet, 2017). It was installed in September 2018 and operated continuously from that point onwards with the exception of several data outages (at times lasting a couple months) due to issues with the instrument's viewing hatch. Measurements are taken continuously and autonomously, except during precipitation when the hatch is closed to prevent damage to the optics, every 47 seconds using seven optical filters: 7.9-9.5, 10-12, 17-18.5, 17.25-19.75, 18.5-20.5, 20.5-22.5, and 22.5-27.5 μm . Due to the strong variation in emissivity between small and large ice crystals in this spectral region, the FIRR bands are sensitive to cloud phase, optical thickness, and microphysical properties. The bands with a wavelength $> 17 \mu\text{m}$ are also very sensitive to small variations in atmospheric water vapour. Detector linearity and radiometric accuracy tests performed at LR Tech and again in the field at Iqaluit confirmed the instrument's accuracy and precision of $<0.1\%$ and $\pm 0.01 \text{ W m}^{-2} \text{ sr}^{-1}$.

250 Each day, the FIRR produces a series of .EEF files totaling 13.8 GB per day. These files are only readable by the licensed EDGAR software (LR Tech Inc.) and contain all unprocessed raw data elements, including housekeeping data. From the .EEF files, EDGAR produces a netCDF file once per day at the end of each day of approximately 1 MB in size containing only essential, processed data (radiance values).

3.1.6 FM-120 Fog Monitor Device

255 The Droplet Measurement Technologies (DMT) FM-120 Fog Monitor Device (FMD) provides continuous and autonomous in-situ observations of PSDs between 2 to 50 μm . It was installed in September 2018 and operated continuously from that point onwards with limited data gaps. By processing the PSD observations, the number concentration, liquid water content, fog intensity, water vapour, and extinction/visibility observations at the surface can be retrieved. As such, its observations are crucial for detecting and understanding the evolution of fog microphysical processes (Gultepe et al., 2017).

260 3.1.7 Precipitation Occurrence Sensor System

The precipitation occurrence sensor system (POSS) is a bistatic X-band Doppler radar designed in-house by ECCC (Sheppard and Joe, 2008; Sheppard et al., 2021). It was installed in September 2018 and operated continuously from that point onwards with limited data gaps. It measures a signal whose frequency is proportional to the raindrop Doppler velocity and whose amplitude is proportional to the raindrop diameter. This provides autonomous and continuous measurements of the precipitation type, rate, raindrop size distribution, and reflectivity. Such high temporal resolution measurements of precipitation are particularly useful for a variety of applications ranging from nowcasting to long-term climatological studies.

3.1.8 Ceilometer

270 The Vaisala CL31 and CL51 ceilometers are lidar instruments that provide aerosol backscatter profile observations to retrieve cloud information such as cloud height, cloud amount (octa, intensity), and aerosol layers up to a range of 7 km a.g.l. A Vaisala CL31 was initially installed at the supersite until September 24 2018; the unit was swapped out for the higher-powered Vaisala CL51 model (operating from September 24 2018 - ongoing). The ceilometers were operated at 5 m vertical resolution and output data in Vaisala's Data Message 2 format (see user manual). Estimates of the PBL height (or mixing layer height) were retrieved from the ceilometer's aerosol backscatter observations using algorithms developed in-house based on existing methodologies (e.g., Kotthaus et al., 2020).

3.1.9 Doppler lidar

280 Two identical scanning Halo Photonics Inc. StreamLine XR Doppler lidars were deployed to Iqaluit; one at the main supersite (operated continuously with limited data gaps) and one at T121 (operated continuously from August 29 2016-onwards but encountered several data outages lasting hours to weeks due to loss of power at T121). The lidars provide accurate observations of aerosol backscatter, depolarization ratio, and Doppler velocity at high temporal- and spatial-resolution along the lidar's beam (radial direction), as well as vertical wind profile observations. While they can scan in all directions like a weather radar,

their observations are limited to the PBL. Their ability to perform rapid scans enables them to observe fast-evolving meteorological features, such as lake breezes, stratified wind layers, low-level jets, and cloud microphysical properties (e.g.,
285 Mariani et al., 2018a; 2018b; Thériault et al., 2021).

Both lidars operated using the same configuration settings and scan strategies as outlined in Mariani et al. (2020a). Vertical staring, over-the-top North-South/East-West/up-valley (135° azimuth) range-height indicator (RHI), plan position indicator (PPI) (4° elevation), Doppler beam swinging, and eight-beam velocity-azimuth display (VAD) vertical wind profile scans were
290 repeated on a 10 minute cycle. The retrieved wind profiles are highly accurate, comparable to radiosonde observations (see Table 1, ‘accuracy’) based on inter-comparison studies (Mariani et al., 2020a). A single raw .hpl file was generated for each scan; each file contains all metadata and measurements during that scan, including scan position (azimuth/elevation), aerosol backscatter, signal to noise ratio, intensity, and Doppler velocity for each range gate. These raw .hpl files were quality-controlled and post-processed processed to produce final wind measurement products (e.g., vertical wind profile) output in
295 standard ASCII files (Mariani et al., 2018a).

3.1.10 DIAL and Raman water vapour lidars

The Vaisala pre-production broadband differential absorption lidar (DIAL) was the first commercial system capable of performing continuous (night and day) observations of the vertical water vapour mass mixing ratio profile. Its design includes
300 two vertically-pointing measurement units placed side-by-side, contained within a larger shelter, with a Vaisala CL-series ceilometer-type telescope (Dabberdt et al., 2016; Roininen et al., 2017). The DIAL underwent initial testing in Helsinki and Toronto before being deployed to Iqaluit (Mariani et al., 2020b). It was installed in September 2018 and operated continuously from that point onwards with limited data gaps, except for a large gap after June 20 2020 when the instrument required maintenance. This new system and its predecessor were extensively evaluated in different climates and demonstrated excellent
305 agreement with independent water vapour profile observations from co-located radiosondes, radiometers, UAV’s, and Raman lidars (Newsom et al., 2020; Mariani et al., 2021; Gaffard et al., 2021).

Water vapour profiles were generated using a 20-min running average up to 3 km a.g.l. (maximum range) output to an ASCII file every minute. Aerosol backscatter profiles were collected every minute up to 14.4 km a.g.l. and output to a separate ASCII
310 file. Estimates of the uncertainty in the water vapor profile and maximum effective range are provided by the DIAL’s quality control algorithm (Newsom et al., 2020). The DIAL’s unique 24 hr continuous water vapour observations enable detailed model inter-comparison studies and measurements of height-resolved diurnal water vapour cycles (Hicks-Jalali et al., 2021). The DIAL stopped collecting observations on June 20 2020 in order to perform repairs; it will be redeployed to the Toronto area.

315

320 A second water vapour lidar, the Canadian Autonomous Arctic Aerosol Lidar (CAAAL), was installed at the Iqaluit supersite. This Raman lidar conducts measurements of the vertical structure of particulate matter, except during precipitation, up to 15 km a.g.l. (Strawbridge et al., 2013; Strawbridge et al., 2018). The lidar was housed in a trailer and was designed and built in-house by ECCC. The lidar conducted simultaneous measurements of aerosol profiles at three wavelengths, including particle size and shape, depolarization ratio measurements at 355 nm, and night-time water vapor mixing ratio measurements using Raman scattering signals at 387 and 407 nm. Its water vapour mixing ratio observations were routinely calibrated using the radiosonde observations at Iqaluit. The CAAAL was redeployed to southern Canada and ended operations at Iqaluit on February 28 2019. Note that the raw CAAAL data is not available on the CAWS archive but can be accessed via coralnet.ca (a password-encoded website that can be accessed by sending a request to Kevin.Strawbridge@ec.gc.ca).

325

3.1.11 Ka-band Radar

330 A Metek Doppler Ka-band weather radar (Bauer-Pfundstein, 2007; Gorsdorf et al., 2015) provides observations of Doppler velocity, backscatter, and depolarization ratio observations at the Iqaluit site. The scanning radar operates at 35.1 GHz using a 30 kW magnetron with a pulse width of 2.0 μ s, a pulse repetition frequency of 5 kHz, and a range resolution of 30 m to a maximum range of 25 km. It was configured to repeat VAD vertical wind profile scans along with several low-elevation PPI and over-the-top RHI scans every 10 minutes, similar to the Doppler lidars. The radar's rapid scan rate enables observation of fast-moving meteorological features such as stratified wind layers and its high sensitivity to light precipitation complements observations from the Doppler lidar (Mariani et al., 2018b). During the study period, the radar operated continuously except for when it experienced several outages caused by inclement weather resulting in a substantial down-time of nearly 50%. Some of these outages lasted several months before repairs could be performed. As such, its available dataset is limited compared to the other instruments at Iqaluit.

335

3.1.12 WXT520 Weather Sensor

340 Thee Vaisala WXT520 weather sensors were deployed at Iqaluit: two at the main supersite (one co-located with the FIRR instrument and one at the top of a 10 m tower near the Doppler lidar), and one beside the Doppler lidar at the T121 site. They all operated continuously with limited data gaps except for the one located at T121 which encountered longer downtime periods due to power outages at T121. Each instrument contains a suite of sensors that conduct in-situ measurements of atmospheric temperature, relative humidity, pressure, wind speed and direction, and precipitation rate (Vaisala, 2012). Data were collected at 1 minute resolution.

345

3.1.13 Other Observations

Additional measurements of precipitation were conducted at the Iqaluit supersite as part of the WMO's Solid Precipitation Inter-Comparison Experiment (SPICE) and the Canadian SPICE (C-SPICE) projects since 2013 (Nitu et al., 2018). The reference configuration used in WMO-SPICE for the measurement of snowfall was the Double Fence Automated Reference (DFAR), as shown in Figure 3. It was designed and characterized for this inter-comparison and employs a suite of instruments, including an automated precipitation gauge (either a Geonor T-200B3 or an OTT Pluvio2) in a single-Alter shield, surrounded by a large octagonal double fence. The DFAR reference measurement incorporates precipitation amount, precipitation occurrence (yes or no) and type (where available) from a sensitive precipitation detector, and environmental conditions such as wind speed and temperature; the reference dataset is therefore a composite dataset from multiple instruments. Additional automatic gauges deployed near the DFAR include single-alter (SA) shielded Geonor weighing gauge (600 mm), unshielded (UN) and SA shielded Pluvio2 weighing gauges (1500 mm), and a Thies Laser Precipitation Monitor (LPM) present weather sensor (not shown in Figure 3). This suite of instruments belong to and are maintained by MSC's Observing Systems and Engineering Section (OSE). Data quality processing of the 6-second Geonor and Pluvio weighing gauges include time formatting, the application of maximum, minimum and data jump filters, manual quality control of the aggregated 1-minute data, and a neutral aggregating filter (Ross et al., 2020). The quality-controlled observations enable detailed study of low snow, cold temperature, high wind and blowing snow conditions. **Observations are available up until September 2018, when the C-SPICE project ended.**

Other observations conducted at the supersite include camera images, aerosol optical depth, and Global Positioning System (GPS) integrated water vapour (IWV). All-sky and 4k pan-tilt camera images were taken every 10 minutes at the site from three vantage points, including multiple sky and cardinal direction views. These images help verify weather conditions at the site and also acted as means to visually check the instruments remotely (e.g., confirm the absence of ice/snow on optics or domes). An aerosol robotic network (AERONET) Cimel sun photometer measured aerosol optical and columnar microphysical properties. Its data is not part of the CAWS archive but it is openly available via <https://aeronet.gsfc.nasa.gov/>. The Iqaluit ground-based GPS receiver is managed by National Resources Canada and the International GNSS Service (IGS). The station was installed in 2009 and has been operating continuously since with less than 2% downtime. The IGS provides IWV and other products at 5 minute resolution, daily, in ASCII format (Jones et al., 2020). Its data is not available as part of this CAWS archive but it is openly accessible via <http://geodesy.unr.edu/> (Blewitt et al., 2018).

3.2 Whitehorse Supersite Instrumentation

A similar but smaller suite of ground-based instruments were deployed to the Whitehorse supersite in 2017 as part of the CAWS project. All instruments were fully automated and operated continuously without an operator at the site. Most of the

380 meteorological parameters observed at the supersite were unique for the region and not observed by the MSC weather station
2.9 km away. Data collection was continuous except for brief power outages when airport runway maintenance was required
(seasonal). A complete list of all instruments and their locations is provided in Table 2; photos of the site and its instruments
are provided in Figure 5. An example of some of the enhanced surface meteorological observations collected from 2018 to
2022 (almost its entire study period) is provided in Figure 6. Due to a planned expansion of the Whitehorse airport facility, the
site was decommissioned in June 2022.

385 As with Iqaluit, the Whitehorse supersite is equipped with an FM-120 FMD, Vaisala CL51 ceilometer, Halo Photonics
Streamline XR Doppler lidar, WXT520 weather sensor, Vaisala FS11P present weather sensor, and all-sky and 4 k pan-tilt
cameras. These instruments were configured and operated in a nearly-identical manner as their counterparts at the Iqaluit
supersite. Thus, the instruments listed in the following subsections are only the ones unique to the Whitehorse supersite. Note
that the Vaisala FS11P present weather sensor deployed to Whitehorse is similar in design and operation to the Vaisala PWD52
390 in Iqaluit (Sect. 3.1.2) operated continuously during the study period up until June 2020, at which point it experienced a laser
failure and had to be removed.

3.2.1 Meteorological Service of Canada Weather Station

The Whitehorse MSC weather station is 2.9 km NW of the Whitehorse supersite; as such it operates completely independently.
395 It has been in operation since 1900. All WMO-standard meteorological surface and upper air (radiosonde; WMO station code
71964) observations are conducted identical to those at the Iqaluit MSC weather office described in Sect. 3.1.1.

3.2.2 Particle size and Precipitation Rate

A shielded OTT Pluvio2 weighing gauge provides surface measurements of precipitation amount and intensity (mm). It
400 operated continuously with limited data gaps. These instruments are commonly used in various climates and have become an
established and reliable measurement platform (Milewska et al., 2019; Smith et al., 2022). No adjustment for wind undercatch
of solid precipitation was performed; as such this dataset should be used with caution.

An optical disdrometer, the OTT Parsivel, provides surface measurements of hydrometeors, including particle size, velocity,
405 and precipitation rate. It operated continuously with limited data gaps. Observations from the Parsivel have been processed to
estimate the precipitation type, intensity, and kinetic energy (Battaglia et al., 2010; Tokay et al., 2014). The instrument uses a
transmitter and receiver separated by a small distance to remotely measure the properties of hydrometeors falling in between
the two sensors.

410 A DMT meteorological particle spectrometer (MPS) is a precipitation-measuring optical disdrometer. It operated continuously with limited data gaps. It measures hydrometeor size distribution and fall velocity, enabling estimates of the precipitation rate for droplets ranging from 50 μm to > 6.4 mm. The instrument processes 2D images of the hydrometeors and a 1D histogram of particle sizes to produce its measurements (Montero-Martinez et al., 2009).

415 3.2.3 X-band Radar

The Selex/Leonardo METEOR 60DX mobile X-band dual-polarization radar with a 2.4 m dish was moved from Vancouver Island, where it previously operated during the Olympic Mountains Experiment (OLYMPEX), to Whitehorse in December 2017 (Hudak et al., 2016; Houze et al., 2017). The radar provides line-of-sight wind speed and direction, cloud & fog backscatter, and depolarization ratio observations, similar to the Ka-band radar at Iqaluit. The scanning radar operates at 9472 MHz using a 75 kW magnetron with a pulse width of 1.0 μs and a dual pulse repetition frequency of 1200/900 Hz. It operated on a five minute cycle conducting three PPI scans (1.5, 3.5, and 5.0° elevation), two RHI scans (170 and 350° azimuth), and a vertical stare. Sector blocking was applied from 200 to 275° in azimuth and below 60° in elevation to protect the other instruments at the supersite, which were located only a few metres away in this direction on a raised platform. Similar to Iqaluit's Ka-band radar, the X-band radar operated continuously during the study period, but experienced several outages caused by inclement weather as well as airport operations requiring the radar to be turned off; as such, its available dataset is limited compared to the other instruments at Whitehorse.

3.2.4 Black Globe Temperature Sensor

A Campbell Scientific black globe temperature sensor provides measurements of heat stress. It uses a thermistor inside a hollow copper sphere that is painted black to measure radiant temperature and operated continuously with limited data gaps. Combined with the measurement of ambient air and wet-bulb temperatures, it is used to calculate the wet-bulb globe temperature (WBGT) index, which is crucial for observing the environmental heat stress felt by an individual.

3.3 Data Storage Rules and Identifiers

435 All geophysical variables observed at the Iqaluit and Whitehorse sites were archived as raw data files and processed in the same manner. Several levels of data processing were published; raw (level 0) data with no quality control (QC) imposed was made available for all instruments, enabling the user to impose their own QC algorithms. As such, all raw data files should be treated with caution, particularly for the radiation flux observations which typically require additional QC processing prior to analysis. Processed data sets (for a limited number of instruments; e.g., lidar VAD wind profiles) are also available as flat text files and images (.pngs). For the processed products, notes in the published readme files point to the type of QC algorithms

applied and whom to contact to obtain processing codes, QC algorithms, or more information in general. In all cases, time is reported as UTC and heights are a.g.l. When no data was available (due to the instrument being down or loss of power at the site), gaps exist or the value -9999 was used. When instruments were maintained and/or recalibrated by technicians visiting the site (roughly twice a year), an identifier in the published metadata is included to mark these service visits.

445

4 Sample of Meteorological Data during High-Impact Weather Events

4.1 Iqaluit blizzard: November 23 2018

Integrated observations were collected during a blizzard on November 23 2018 at the Iqaluit supersite. A low pressure system brought thick low-level clouds (1 to 2 km a.g.l.) to Iqaluit on the evening of November 22 2018. Strong surface winds > 15 m/s produced a mixture of drifting and blowing snow into the next day. Combined with precipitating snow from the low-level clouds, surface visibility conditions were near-zero for the entire morning of November 23rd, as observed by the PWD52 and the METAR reports. Surface temperatures ranged from -19.9 to -23.7 °C throughout the blizzard, as measured by the WXT520. The storm lasted until around 11:00 UTC, when surface winds decreased to < 13 m/s and precipitation ended. This improved surface visibility to > 5 km. Thick persistent cloud cover remained after the storm, including thin ice clouds and ice crystals (diamond dust). A total of 1.8 mm of precipitation was measured at the MSC weather station during the blizzard.

Figures 7 and 8 provide examples of some of the data collected at the supersite during the blizzard. Aerosol backscatter observations from the CL51 ceilometer (including meteorological type classifications), water vapour profile observations from the DIAL, downwelling brightness temperatures from the FIRR, and short- and longwave fluxes measured by the radiation flux sensor suite are provided in Figure 7. These observations characterize the radiative budget at the site during and after the blizzard with large variations observed as a result of the changing snow and cloud conditions. For instance, a period of clear skies between 14:45 and 16:45 UTC shown in Fig. 7(a) resulted in a sudden decrease in the downwelling brightness temperatures in Fig. 7(c). The extremely low water vapour profile concentrations (Fig. 7(b), ~0.5 g/kg) enable the FIRR channels that are mostly transparent to other trace gases (e.g., 10-12 μm) to effectively observe the downwelling radiation from near-space. Changes in the short- and longwave fluxes (Fig. 7d-e) also corresponded with changing cloud cover; note they are only shown up to 14:00 UTC due to an instrument malfunction recorded soon after the blizzard ended.

Observations of microphysical particle properties during the blizzard are provided in Figure 8. The number concentration, liquid water concentration, and particle size as measured by the FMD are provided in Fig. 8(a-b), precipitation type (blue) and rate (green) observations from the PWD52 are shown in Fig. 8(c), and a PIP particle image of snowflakes taken at 08:37 UTC is shown in Fig. 8(d) with corresponding DSDs and fall velocities processed by the PIP shown in Figs. 8(e, f). Note that Figs. 8(e, f) are shown only up to 12:00 UTC since there was no blowing or precipitating snow recorded after this time. All

observations are consistent with the presence of moderate or heavy blowing and/or precipitating snow, ranging in sizes mostly < 5 mm (some cases up to 10 mm). Residual blowing snow was observed by the FMD and PIP for several hours after the main snowfall event (10:00 UTC onwards).

4.2.0 Whitehorse blizzard: Dec 16-17 2019

A HIW event occurred at the Erik Nielsen Whitehorse International Airport on December 16-17 2019. A low pressure system brought persistent thick and precipitating low-level clouds (~400 m a.g.l.) on December 16 that covered the entire Whitehorse valley, including the airport. Surface temperatures ranged between -15.4 and -11.5 °C, as observed by the WXT520. A total of 1.6 mm of precipitation was measured at the MSC weather station during the blizzard.

Periods of near-zero surface visibility, precipitating wet snow, and a very low cloud base made flights into and out of Whitehorse hazardous, particularly given the complex mountainous terrain surrounding the airport. Several flights were cancelled or diverted, severely impacting transportation for the entire Western Arctic region. One two-hour flight, Air Canada AC 279, which departed from Vancouver, British Columbia, to Whitehorse on Dec 16, received notable media attention when it turned into a two-day international trip, detouring to Anchorage, Alaska after it was deemed unsafe to land in Whitehorse due to the inclement weather.

Data collected by the PWD52 at the Whitehorse supersite during the event are shown in Figure 9, including the luminance, precipitation type (blue) and rate (grey), and surface visibility. Periods of near-zero surface visibility are correlated with increased precipitating snow rates. Luminance was zero for most of the day (except when clouds began to dissipate after 18:00 UTC), further exacerbating the poor visibility conditions. Periods of intermittent precipitation continued with intermittent thick cloud cover throughout the evening and on December 17.

Doppler lidar and X-band radar remote sensing observations provided high temporal and spatial resolution upper air observations during the HIW event. Such observations are extremely relevant for aviation nowcasting operations. Doppler lidar vertical profiles of the attenuated backscatter, vertical velocity (w -component of the wind), depolarization ratio (to distinguish ice/water composition), and horizontal winds (u - and v -components of the wind) on December 16 2019 are shown in Figs. 10(a-d), respectively. A low cloud ceiling of ~400 m a.g.l. was relatively constant throughout most of the day, limiting the lidar's vertical observations to this altitude. Vertical velocities indicated mixing and turbulent motions within the PBL throughout most of the day, with strong horizontal winds (~15 m/s) observed at the cloud base (~500 m a.g.l.) after 16:00 UTC and a strong wind shear at ~1.25 km a.g.l. Note that these wind conditions occurred within the Whitehorse valley, below the mountain peaks, producing dangerous circumstances for aircraft operations, particularly considering the near-zero visibility.

The X-band radar's 3.5° elevation PPI scans at 9:45 UTC provide snapshots of the horizontal extent of the storm during its peak in precipitation throughout the Whitehorse valley. Radar reflectivity is shown in Fig. 10(e) and differential reflectivity (Z_{DR}) is shown in Fig. 10(f). Both the Doppler lidar depolarization ratio vertical profiles (Fig. 10c) and the X-band radar's horizontal differential reflectivity (Fig. 10f) indicate the presence of wet precipitating snow throughout and above the entire valley. Note that the X-band data in Figs. 10(e-f) is limited not by its range but due to the topography of the region, with nearby mountains blocking the radar's beam. Though logistically challenging, if the radar was moved to a higher position near the top of the nearby mountains the improved sightlines would significantly improve its ability to detect such storm systems earlier.

5 Data Availability

The CAWS dataset is available via the Government of Canada Open Data Portal and can be accessed at: <https://doi.org/10.18164/ff771396-b22c-4bc3-844d-38fc697049e9> (Iqaluit supersite, Mariani et al., 2022a) and <https://doi.org/10.18164/d92ed3cf-4ba0-4473-beec-357ec45b0e78> (Whitehorse supersite, Mariani et al., 2022b). Meteorological Service of Canada surface and radiosonde data is available via weather.gc.ca.

6 Final Remarks

Two unique datasets of enhanced meteorological observations were collected at the Iqaluit and Whitehorse supersites during the CAWS project. Data at Iqaluit were collected starting September 2015 and is partially ongoing (depending on the instrument) while data collected at Whitehorse was collected from November 2017 to June 2022. These new datasets fill large gaps in meteorological observations in the Arctic and improve existing observing capabilities via the deployment and integration of new, at times prototype remote sensing technologies. This data is being used to fulfill the CAWS project's goal of providing recommendations on a cost-effective Arctic observing system; improved Arctic infrastructure, satellite calibration/validation, new weather products for Northern communities, and enhanced near-real time HIW observing capabilities have also been achieved as a result of CAWS.

The combination of multiple in-situ and remote sensing instruments deployed at the two sites provides an unprecedented wealth of integrated meteorological observations for the Canadian Arctic. The sites' automated and continuous observations of vertically-resolved winds, water vapour, clouds and aerosols, visibility, radiation fluxes, and precipitation are a unique high-resolution dataset encompassing all essential meteorological parameters from the sub-surface soil up to PBL and beyond. Overall, the CAWS dataset will significantly contribute to our understanding of synoptic and fine-scale meteorological

535 processes in the Arctic, including cloud microphysical processes, the radiative budget, HIW detection and prediction, nowcasting, PBL dynamics, and NWP verification, assimilation, and processes, particularly in the context of WMO YOPP.

Author contributions

540 ZM wrote the first draft of the manuscript and conducted analyses. ZM, SHJ, EM, LP, and PR conducted scientific analyses and created plots. ZM, LH, and RC managed data archiving and publication to the Government of Canada Open Data Portal. All contributed to the writing and the editing of the manuscript.

Competing interests

545 The authors declare that they have no conflict of interest.

Disclaimer

Use of specific instrument manufacturers/models and suppliers mentioned in the manuscript and/or used at the supersites is not a commercial endorsement of their products.

550

Acknowledgements

Special thanks to Sorin Pinzariu, Michael Harwood, Robert Reed, Reno Sit, Jason Iwachow, Michael Travis, Bernard Firanski, and Daniel Coulombe (ECCC) for their help with instrumentation at the Iqaluit site. Thank you to Larry Bliven and David Wolff (NASA) for providing the PIP instrument and its data processing package. Thank you to the Meteorological Service of Canada radiosonde operators and the Observing System and Engineering Division's management of the Iqaluit MSC weather station. Thank you to the Erik Nielsen Whitehorse International Airport Management and the Government of Yukon for providing access to the Whitehorse Airport field in order to commission the Whitehorse supersite. All data products are produced by ECCC and are available via obsr.ca, the Government of Canada Open Data Portal (<https://doi.org/10.18164/ff771396-b22c-4bc3-844d-38fc697049e9> and <https://doi.org/10.18164/d92ed3cf-4ba0-4473-beec-357ec45b0e78>), or upon request.

555
560

References

- Arctic Council (2017): Telecommunications Infrastructure in the Arctic: a circumpolar assessment, accessed 3 Feb 2019, <https://oaarchive.arctic-council.org/handle/11374/1924>.
- 565 Barthazy, E., and R. Schefold (2006): Fall velocity of snowflakes of different riming degree and crystal types. *Atmos. Res.*, 82, 391-398.
- Battaglia, A., E. Rustemeier, A. Tokay, U. Blahak, and C. Simmer (2010): Parsivel snow observations: a critical assessment. *J. Atmos. Ocean. Tech.*, 27, 333-344, DOI: 10.1175/2009JTECHA1332.1.
- Bauer-Pfundstein, M.R. (2007): Target Separation and Classification using Cloud Radar Doppler-Spectra. In Proceedings of
570 the AMS 33rd International Conference on Radar Meteorology, Cairns, Australia, 6–10 August 2007.
- Blewitt, G., W. C. Hammond, and C. Kreemer (2018): Harnessing the GPS data explosion for interdisciplinary science, *EOS*, 99, <https://doi.org/10.1029/2018EO104623>.
- Cassano, J., Higgins, M., and Seefeldt, M. (2011): Performance of the Weather Research and Forecasting Model for Month-
Long Pan-Arctic Simulations. *Mon. Wea. Rev.*; 139, 3469–3488.
- 575 Chou, C.C., P.J. Kushner, S. Laroche, Z. Mariani, S. Melo, and C.G. Fletcher (2021): Validation of the Aeolus Level-2B wind product over Northern Canada and the Arctic. *Atmos. Meas. Tech.*, submitted August 20 2021.
- Cober, S. G., Isaac, G. A., and Korolev, A. V. (2001): Assessing the Rosemount Icing Detector with In Situ Measurements, *Journal of Atmospheric and Oceanic Technology*, 18(4), 515-528.
- Cox, C. J., Morris, S. M., Uttal, T., Burgener, R., Hall, E., Kutchenreiter, M., McComiskey, A., Long, C., Thomas, B., and
580 Wendell, J. (2021): The De-Icing Comparison Experiment (D-ICE): a study of broadband radiometric measurements under icing conditions in the Arctic. *Atmos. Meas. Tech.*, 14(2), 1205-1224. <https://doi.org/10.5194/amt-14-1205-2021>.
- Dabberdt, W.F.; C. Munkel, J. Kallio, M. Komppula, S. Laukkanen, and E.J. O'Connor (2016): Advances in Continuous Atmospheric Boundary Layer Humidity Profiling with a Compact DIAL Instrument. In Proceedings of the 18th Symposium on Meteorological Observation and Instrumentation, New Orleans, LA, USA, 13 January 2016; American
585 Meteorological Society: Boston, MA, USA, 2016; Volume 8.4. Available online: <https://ams.confex.com/ams/96Annual/webprogram/Paper285586.html> (accessed on 14 July 2021).

- Ford, J.D., G. McDowell, J. Shirley, M. Pitre, R. Siewierski, W. Gough, F. Duerden, et al. (2013): The dynamic multiscale nature of climate change vulnerability: An Inuit harvesting example. *Annals of the Association of American Geographers*, 103(5):1193-1211.
- 590 Gaffard, C., Z. Li, D. Harrison, R. Lehtinen, and R. Roininen (2021): Evaluation of a Prototype Broadband Water-Vapour Profiling Differential Absorption Lidar at Cardington, UK. *Atmosphere*, 12, 1521. <https://doi.org/10.3390/atmos12111521>.
- Görsdorf, Ulrich, V. Lehmann, M. Bauer-Pfundstein, G. Peters, D. Vavriv, V. Vinogradov, and V. Volkov (2015): A 35-GHz Polarimetric Doppler Radar for Long-Term Observations of Cloud Parameters—Description of System and Data Processing. *Journal of Atmospheric and Oceanic Technology*, 32.4: 675-690. [https://doi.org/10.1175/JTECH-D-14-](https://doi.org/10.1175/JTECH-D-14-00066.1)
- 595 [00066.1](https://doi.org/10.1175/JTECH-D-14-00066.1)
- Government of Canada (2016): Quadrennial Search and Rescue Review. 16pp.
- GRAW (2020): Available online: <https://www.graw.de/products/radiosondes/dfm-09/> (accessed on 19 January 2022).
- Gultepe, I., R. Rabin, R. Ware, and M. Pavolonis (2016): Chapter three – light snow precipitation and effects on weather and climate. *Advances in Geophysics*, 57, 147-210, <https://doi.org/10.1016/bs.agph.2016.09.001>.
- 600 Gultepe, I., A. Heymsfield, P. Field, and D. Axisa (2017): Ice-phase precipitation. *AMS Meteorological Monographs*, 58, 6.1-6.36, <https://doi.org/10.1175/AMSMONOGRAPHS-D-16-0013.1>.
- Gunn, R. and Kinzer, G. D. (1949): The terminal velocity of fall for water droplets in stagnant air, *Journal of Meteorology*, 6, 243–248, [https://doi.org/10.1175/1520-0469\(1949\)006<0243:TTVOFT>2.0.CO;2](https://doi.org/10.1175/1520-0469(1949)006<0243:TTVOFT>2.0.CO;2).**
- Hicks–Jalali, S., Mariani, Z. & Crawford, R.W. (2021): DIAL and GNSS observations of the diurnal water-vapour cycle above
- 605 Iqaluit, Nunavut. *Q J R Meteorol Soc*, 147 (741), 4228– 4250. <https://doi.org/10.1002/qj.4175>.
- Houze, R. A., Jr., McMurdie, L. A., Petersen, W. A., Schwaller, M. R., Baccus, W., Lundquist, J. D., Mass, C. F., Nijssen, B., Rutledge, S. A., Hudak, D. R., Tanelli, S., Mace, G. G., Poellot, M. R., Lettenmaier, D. P., Zagrodnik, J. P., Rowe, A. K., DeHart, J. C., Madaus, L. E., Barnes, H. C., and Chandrasekar, V. (2017): The Olympic Mountains Experiment (OLYMPEX), *Bulletin of the American Meteorological Society*, 98(10), 2167-2188. Retrieved Mar 21, 2022, from
- 610 <https://journals.ametsoc.org/view/journals/bams/98/10/bams-d-16-0182.1.xml>

- Hudak, D., P. Rodriguez, N. Donaldson, and D. Kirshbaum (2016): OLYMPEX Canada. In Proceedings of the AMS 17th Conference on Mountain Meteorology, Burlington, Vermont, U.S.A., 27 June - 1 July 2016. <https://ams.confex.com/ams/17Mountain/webprogram/Paper296448.html>.
- 615 Joe, P., Melo, S., Burrows, W. R., Casati, B., Crawford, R. W., Dehghan, A., Gascon, G., Mariani, Z., Milbrandt, J., and Strawbridge, K. (2020): The Canadian Arctic Weather Science Project: Introduction to the Iqaluit Site, Bull. Amer. Meteor. Soc. (Open Access), 101(2), E109-E128
- Illingworth, A.J., D. Cimini, C. Gaard, M. Haeelin, V. Lehmann, U. Löhnert, E.J. O'Connor, and D. Rieux (2015): Exploiting existing ground-based remote sensing networks to improve high-resolution weather forecasts. Bull. Am. Meteorol. Soc. 2015, 96, 2107–2125
- 620 Jones, J., Guerova, G., Dousa, J., Dick, G., de Haan, S., Pottiaux, E., Bock, O., Pacione, R., and van Malderen, R., (2020): Advanced GNSS Tropospheric Products for Monitoring Severe Weather Events and Climate. Springer Nature Switzerland, Cham, Switzerland.
- Koltzow, M.; B. Casati, E. Bazile, T. Haiden, and T. Valkonen (2019): An NWP model intercomparison of surface weather parameters in the European Arctic during the year of polar prediction special observing period northern hemisphere 1. 625 Weather Forecast. 2019, 34, 959–983.
- Kotthaus, S., M. Haeffelin, M.-A. Drouin, J.-C. Dupont, S. Grimmond, A. Haeefele, M. Hervo, Y. Poltera, and M. Wiegner (2020): Tailored Algorithms for the Detection of the Atmospheric Boundary Layer Height from Common Automatic Lidars and Ceilometers (ALC). Remote Sens. 2020, 12, 3259. <https://doi.org/10.3390/rs12193259>.
- Kumer, V.M.; J. Reuder, and B.R. Furevik (2014): A comparison of LiDAR and radiosonde wind measurements. Energy 630 Procedia 2014, 53, 214–220.
- Lesins, G., L. Bourdages, T.J. Duck, J.R. Drummond, E. Eloranta, and V.P. Walden (2009): Large surface radiative forcing from topographic blowing snow residuals measured in the High Arctic at Eureka. Atmos. Chem. Phys., 9, 1847-1862.
- Libois, Q., C. Proulx, L. Ivanescu, L. Coursol, L. S. Pelletier, Y. Bouzid, F. Barbero, Éric Girard, and J.-P. Blanchet (2016): A microbolometer-based far infrared radiometer to study thin ice clouds in the Arctic. Atmos. Meas. Tech., 9, 1817–1832, 635 <https://doi.org/10.5194/amt-9-1817-2016>.

- Libois, Q., and J.-P. Blanchet (2017): Added value of far-infrared radiometry for remote sensing of ice clouds, *J. Geophys. Res. Atmos.*, 122, 6541–6564, doi:10.1002/2016JD026423.
- Liu, X.C., T.C. Gao, and L. Liu (2013): a comparison of rainfall measurements from multiple instruments. *Atmos. Meas. Tech.*, 6, 1585-1595, doi:10.5194/amt-6-1585-2013.
- 640 Mariani, Z.; Dehghan, A.; Sills, D.M.; Joe, P (2018a): Observations of Lake Breeze Events during the Toronto 2015 Pan-American Games. *Bound. Layer Meteorol.* 2018, 166, 113–135, doi:10.1007/s10546-017-0289-3.
- Mariani, Z., A. Dehghan, G. Gascon, P. Joe, D. Hudak, K. Strawbridge, and J. Corriveau (2018b): Multi-instrument observations of prolonged stratified wind layers at Iqaluit, Nunavut. *Geophys. Res. Lett.* 2018, 45, 1654–1660.
- Mariani, Z.; R. Crawford, B. Casati, and F. Lemay (2020a): A Multi-Year Evaluation of Doppler Lidar Wind-Profile
645 Observations in the Arctic. *Remote Sens.* 2020, 12, 323. <https://doi.org/10.3390/rs12020323>
- Mariani, Z.; N. Stanton, J. Whiteway, and R. Lehtinen (2020b): Toronto Water Vapor Lidar Inter-Comparison Campaign. *Remote Sens.* 2020, 12, 3165. <https://doi.org/10.3390/rs12193165>
- Mariani, Z.; S. Hicks-Jalali, K. Strawbridge, J. Gwozdecky, R.W. Crawford, B. Casati, F. Lemay, R. Lehtinen, and P. Tuominen (2021): Evaluation of Arctic Water Vapor Profile Observations from a Differential Absorption Lidar. *Remote
650 Sens.* 2021, 13, 551. <https://doi.org/10.3390/rs13040551>
- Mariani, Z., Huang, L., and R. Crawford (2022a): Iqaluit, Nunavut (CYFB) Enhanced Site for Canadian Arctic Weather Science (CAWS) Project. Government of Canada Open Data Portal, <https://doi.org/10.18164/ff771396-b22c-4bc3-844d-38fc697049e9>.
- Mariani, Z., Huang, L., and R. Crawford (2022b): Whitehorse, Yukon (CYXY) Enhanced Site for Canadian Arctic Weather
655 Science (CAWS) Project. Government of Canada Open Data Portal, <https://doi.org/10.18164/d92ed3cf-4ba0-4473-beec-357ec45b0e78>.
- Milewska, E. J., Vincent, L. A., Hartwell, M. M., Charlesworth, K., and Mekis, É. (2019): Adjusting precipitation amounts from Geonor and Pluvio automated weighing gauges to preserve continuity of observations in Canada, *Can. Water Res. J.*, 44, 127–145, <https://doi.org/10.1080/07011784.2018.1530611>.

- 660 Montero-Martinez, G., Kostinski, A. B., Shaw, R. A., and Garcia-Garcia, F. (2009): Do all raindrops fall at terminal speed?,
Geophys. Res. Lett., 36, L11818, <https://doi.org/10.1029/2008GL037111>.
- Newman, A. J., P. A. Kucera, and L. F. Bliven (2009): Presenting the snowflake video imager (SVI). J. Atmos. Oceanic
Technol., 26, 167–179, <https://doi.org/10.1175/2008JTECHA1148.1>.
- Newsom, R.K.; D.D. Turner, R. Lehtinen, C. Munkel, J. Kallio, and R. Roininen (2020): Evaluation of a Compact Broadband
665 Differential Absorption Lidar for Routine Water Vapor Profiling in the Atmospheric Boundary layer. J. Atmos. Ocean.
Technol. 2020, 37, 47–65.
- Nitu, R., Roulet, Y., Wolff, M., Earle, M., Reverdin, A., Smith, C., Kochendorfer, J., Morin, S., Rasmussen, R., Wong, K.,
Alastru'e, J., Arnold, L., Baker, B., Buisan, S., Collado, J. L., Colli, M., Collins, B., Gaydos, A., Hannula, H.R., Hoover,
J., Joe, P., Kontu, A., Laine, T., Lanza, L., Lanzinger, E., Lee, G.W., Lejeune, Y., Leppanen, L., Mekis, E., Panel, J.,
670 Poikonen, A., Ryu, S., Sabatini, F., Theriault, J., Yang, D., Genthon, C., van den Heuvel, F., Hirasawa, N., Konishi, H.,
Nishimura, K., and Senese, A., (2018): WMO solid precipitation intercomparison experiment (SPICE) (2012–2015).
https://library.wmo.int/doc_num.php?explnum_id=5686.
- Paschke, E.; R. Leinweber, and V. Lehmann (2015): An assessment of the performance of a 1.5 m Doppler lidar for operational
vertical wind profiling based on a 1-year trial. Atmos. Meas. Tech. 2015, 8, 2251–2266.
- 675 Pettersen, C.; Bliven, L.F.; von Lerber, A.; Wood, N.B.; Kulie, M.S.; Mateling, M.E.; Moisseev, D.N.; Munchak, S.J.;
Petersen, W.A.; and Wolff, D.B (2020): The Precipitation Imaging Package: Assessment of Microphysical and Bulk
Characteristics of Snow. Atmosphere 2020, 11, 785. doi: <https://doi.org/10.3390/atmos11080785>
- Pettersen C., Bliven L. F., M. S. Kulie, N. B. Wood, J. Shates, J. Anderson, M. Mateling, W. A. Petersen, A. V. Lerber., D. B.
Wolff (2021): The Precipitation Imaging Package: Phase Partitioning Capabilities. Atmosphere, 2021. Doi:
680 <https://doi.org/10.3390/rs13112183>.
- Riishojgaard, L. (2015): Wind Measurements in the WMO Global Observing System. *ADM-Aeolus Science and Cal/Val
Workshop*; ESRIN, Frascati, 10-13 Feb.
- Roininen, R. and C. Munkel (2017): Results from Continuous Atmospheric Boundary Layer Humidity Profiling with a
Compact DIAL Instrument. In Proceedings of the Eighth Symposium on Lidar Atmospheric Applications, Seattle, WA,

685 USA, 23 January 2017; American Meteorologic Society: Boston, MA, USA, 2017; Volume 12.3. Available online:
<https://ams.confex.com/ams/97Annual/webprogram/Paper301717.html> (accessed on 14 July 2021).

Ross, A., C.D. Smith, and A. Barr (2020): An improved post-processing technique for automatic precipitation gauge time series. *Atmos. Meas. Tech.*, 13, 2979–2994, 2020. <https://doi.org/10.5194/amt-13-2979-2020>.

Schyberg, H. and Randriamampianina, R. (2015): MET Norway plans for contribution to calibration-validation and use of
690 Aeolus winds. *ADM-Aeolus Science and Cal/Val Workshop*; ESRIN, Frascati, 10-13 Feb.

Sheppard, B.E., and P. I. Joe (2008): Performance of the Precipitation Occurrence Sensor System as a Precipitation Gauge. *J. Atmos. Ocean. Tech.*, 25, 196-212, <https://doi.org/10.1175/2007JTECHA957.1>.

Sheppard, B. E., M. Thurai, P. Rodriguez, P. C. Kennedy, and D. Hudak (2021): Improved Precipitation Typing Using POSS Spectral Modal Analysis, *Journal of Atmospheric and Oceanic Technology*, 38(3), 537-554.
695 <https://doi.org/10.1175/JTECH-D-20-0075.1>.

Smith, L. C. and S. R. Stephenson, 2013; New Trans-Arctic Shipping routes navigable by mid century, *Proceedings of the National Academy of Sciences*, 110, 4871-4872.

Smith, C. D., Mekis, E., Hartwell, M., and Ross, A.: The hourly wind-bias adjusted precipitation data set from the Environment and Climate Change Canada automated surface observation network (2001–2019), *Earth Syst. Sci. Data Discuss.* [preprint],
700 <https://doi.org/10.5194/essd-2022-208> , in review, 2022. Data set (assets tab): <https://doi.org/10.18164/6b90d130-4e73-422a-9374-07a2437d7e52>

Statistics Canada (2016): Table 051-0005 - Estimates of population, Canada, provinces and territories, quarterly (persons), CANSIM (database).

Strawbridge, K.B (2013): Developing a portable, autonomous aerosol backscatter lidar for network or remote operations.
705 *Atmos. Meas. Tech.* 2013, 6, 801–816.

Strawbridge, K.B.; Travis, M.S.; Firanski, B.J.; Brook, J.R.; Staebler, R.; and Leblanc, T (2018): A fully autonomous ozone, aerosol and nighttime water vapor lidar: A synergistic approach to profiling the atmosphere in the Canadian oil sands region. *Atmos. Meas. Tech.* 2018, 11, 6735–6759.

- 710 Tiira, J., D. N. Moisseev, A. von Lerber, D. Ori, A. Tokay, L. F. Bliven, and W. Petersen (2016): Ensemble mean density and its connection to other microphysical properties of falling snow as observed in southern Finland. *Atmos. Meas. Tech.*, 9, 4825–4841, <https://doi.org/10.5194/amt-9-4825-2016>.
- Tokay, A., D. Wolff, and W. Petersen (2014): Evaluation of the new version of the laser-optical disdrometer, OTT Parsivel2. *J. Atmos. Ocean. Tech.*, 31, 1276-1288, DOI: 10.1175/JTECH-D-13-00174.1.
- 715 Thériault, J. M., Déry, S. J., Pomeroy, J. W., Smith, H. M., Almonte, J., Bertoincini, A., Crawford, R. W., Desroches-Lapointe, A., Lachapelle, M., Mariani, Z., Mitchell, S., Morris, J. E., Hébert-Pinard, C., Rodriguez, P., and Thompson, H. D. (2021): Meteorological observations collected during the Storms and Precipitation Across the continental Divide Experiment (SPADE), April–June 2019, *Earth Syst. Sci. Data*, 13, 1233–1249, <https://doi.org/10.5194/essd-13-1233-2021>, 2021.
- Vaisala (2007): Vaisala Radiosonde RS92 Measurement Accuracy; Technical Report; Vaisala: Vantaa, Finland, 2007.
- Vaisala (2012): User’s Guide: VaisalaWeather Transmitter WXT520, available at:
720 <https://www.vaisala.com/sites/default/files/documents/M210906EN-C.pdf> (last access: 22 October 2019), 2012.
- von Lerber, A., D. Moisseev, D. A. Marks, W. Petersen, A. M. Harri, and V. Chandrasekar (2018): Validation of GMI snowfall observations by using a combination of weather radar and surface measurements. *J. Appl. Meteor. Climatol.*, 57, 797–820, <https://doi.org/10.1175/JAMC-D-17-0176.1>.
- 725 WMO (2011): Weather extremes in a changing climate: hindsight on foresight. World Meteorological Organization (WMO); WMO-No. 1075, pp. 1-20.
- WMO (2017): Dawson, J., Hoke, W., Lamers, M., Liggett, D., Ljubicic, G., Mills, B., Stewart, E., et al. 2017. Navigating Weather, Water, Ice and Climate Information for Safe Polar Mobilities. WWRP/PPP No.5 – 2017. World Meteorological Organization, Geneva.
- Zhang, L., L. Zhao, C. Xie, G. Liu, L. Gao, Y. Xiao, J. Shi, and Y. Qiao (2015): Intercomparison of Solid Precipitation Derived
730 from the Weighting Rain Gauge and Optical Instruments in the Interior Qinghai-Tibetan Plateau. *Advances in Meteorology*, 2015, Article ID 936724, 11 pp., doi:10.1155/2015/936724.

735 **Table 1:** List of instruments at the Iqaluit supersite including their technical specifications such as instrument manufacturer, date range of observations, meteorological measurement products, and accuracy (where applicable). Not listed: cameras, MSC standard meteorological surface observations, radiosonde observations, and the WMO C-SPICE precipitation sensor test-field instruments. Variables and accuracies reported in the manufacturer’s manual are provided for most instruments; lidar variables and accuracies are from Mariani et al. (2020a, 2021) and Rosemount accuracies are from Cober et al. (2001).
 740 The * denotes an additional instrument of identical design deployed at the T121 Ridge site.

Instrument	Manufacturer	Date Range (YYYYMM) of Observations	Operating Principle	Measurement(s)	Temporal / Geographic Resolution	Accuracy
Precipitation Imaging Package (PIP)	NASA/Wallops	20140911 – ongoing	380 frames per second grey-scale camera with back-lighting	Particle imagery, DSD, precip. rate and density estimation	< 1 min / surface obs.	N/A
Ka-Band Radar	Metek	20150929 – 20191210	Scanning pulsed dual-polarization Doppler radar. 1.2 m dish (0.5° beamwidth).	Line-of-sight wind speed and direction, cloud & fog backscatter, depolarization ratio	10 min / 30 m res. up to ~25 km range	Antenna gain: 50.4 dBi Sensitivity: -52.4 dBZ at 5 km Nyquist velocity: ± 10.7 m/s Velocity resolution: ≥ 0.08 m/s
Ceilometer CL31 / CL51	Vaisala	CL31: 20150929 – 20180924 CL51: 20180924 - ongoing	Pulsed (6.5 kHz) diode laser lidar	Cloud intensity, cloud octa and height, aerosol profiles, PBL height	< 1 min / 5 m vert res. up to 15 km a.g.l.	Distance: better than ±5 m
Present Weather Detector PWD52	Vaisala	20150929 – ongoing	Forward-scatter measurement	Visibility, luminance, precipitation rate and type	< 1 min / surface obs.	Visibility: ±10% up to 10 km Precipitation sensitivity: 0.05 mm/hr
Streamline XR Doppler Lidar (x2*)	Halo Photonics	20150929 – ongoing	Pulsed (10 kHz) scanning at 1.5 µm (Mie scattering)	Line-of-sight wind speed and direction, aerosol backscatter, depolarization ratio	5 min / 3 m res. up to 10 km range (backscatter) or ~2-4 km (Doppler velocity)	Doppler velocity < 0.3 m/s Average vertical wind-profile bias to radiosonde: 0.27 m/s
Rosemount icing detector	Rosemount Engineering	20150929 – ongoing	Magnetostrictive oscillation probe with a sensing cylinder	Detects presence of ice, frost	< 1 min / surface obs.	LWC threshold: 0.007 ±0.010 g/m ³
Weather Sensor WXT520 (x3*)	Vaisala	20150929 – ongoing	Several sensors & transducers housed in a single unit	2.5 m (x2*) and 10 m a.g.l. winds, P, T, RH, and precipitation rate and accumulation	< 1 min / surface obs (a second unit is deployed at 10 m a.g.l.).	Wind speed: ±3% at 10 m/s Wind direction: ± 3° P: ±1hPa (±0.5 hPa 0 to 30 °C), T: ±0.3 °C (at 20 °C) RH: ±3% (0% to 90% RH), ±5% (90% to 100% RH)

						Accum. precip.: < 5%
Canadian Autonomous Arctic Aerosol Lidar (CAAAL)	In-house (ECCC)	20161127 - 20190228	355/532/1064 nm transmitter & 6 channel receiver	Aerosol and water vapour profiles; depolarization ratio	< 1 min / 3.75 m res. up to ~15 km a.g.l. (30 m up to 10 km for water vapour)	Profile uncertainty is provided for each measurement since it varies
FM-120 Fog Monitor Device (FMD)	Droplet Measurement Technologies (DMT)	20180914 – ongoing	Single-particle forward light scattering	Surface particle diameter, number concentration, LWC, Eff. Diam.	< 1 min / surface obs.	Droplets between 2 – 50 μm
Far-IR Radiometer (FIRR)	LR Tech.	20180914 – ongoing	Zenith/Nadir-viewing infrared radiometer	Downwelling IR radiation and brightness temperatures at 7 spectral channels, cloud microphysics	2 min / N/A	Radiometric accuracy: $\pm 0.01 \text{ W/m}^2\text{sr}^{-1}$
Surface radiation flux sensor suite	Campbell Scientific	20180914 – ongoing	Surface radiation pyranometer and pyrgeometers (diffuse and direct)	Up- and downward Shortwave (pyranometer) and up/down/N/E/S/W longwave (pyrgeometers) radiation flux sensors	1 min / N/A	Pyranometer sensitivity: 7-14 $\mu\text{V/W/m}^2$ Pyranometer offset: < 7 W/m^2 Pyrgeometer sensitivity: 5-15 $\mu\text{V/W/m}^2$ Pyrgeometer window heating offset: < 4 W/m^2 Pyrgeometer offset: < 2 W/m^2 Temperature dependence of sensitivity (both): < 1%
DIAL water vapour Lidar	Vaisala	20180914 – 20200620	Pulsed DIAL lidar system	Profiles of aerosol backscatter and water vapour	1 min (20 for water vapour) / 5 m up to 14.4 km (~3 km for water vapour) a.g.l.	Profile uncertainty is provided for each measurement since it varies. Average bias to radiosonde: +0.13 g/kg
Soil probe and SR50ATH snow depth sensor	Campbell Scientific	20180914 – ongoing	Soil probe and ultrasonic distance sensors	Soil water volume and temperature, snow depth	< 1 min / surface obs.	Soil water volume accuracy: $\pm 3\%$ Soil temperature accuracy: $\pm 0.5 \text{ }^\circ\text{C}$ Snow depth: $\pm 1 \text{ cm}$
Precipitation occurrence observation system	In-house (ECCC)	20180914 - ongoing	Meteorological radar profiler for Doppler spectra of hydrometeors	Precipitation type, rate, and backscatter	< 1 min / surface obs.	Doppler frequency resolution: 0.23 m/s Minimum precip. rate: 0.1 mm/hr

745 **Table 2:** Same as Table 1, except for the Whitehorse supersite

Instrument	Manufacturer	Date Range (YYYYMM) of Observations	Operating Principle	Measurement(s)	Temporal / Geographic Resolution	Accuracy
X-Band Radar	Selex/Leonardo	20171215 - 20200201	Scanning pulsed dual-polarization Doppler Radar. 2.4 m dish (0.98° beamwidth).	Cloud backscatter, winds, and precipitation surveillance	5 min / 250 m res. up to ~100 km range	Antenna gain: 44.5 dBi Sensitivity: -1.5 dBZ at 50 km Nyquist velocity: ± 28.5 m/s Velocity resolution: 0.224 m/s
Streamline XR+ Doppler Lidar	Halo Photonics	20171129 - 20220601	Pulsed (10 kHz) scanning at 1.5 μm (Mie scattering)	Line-of-sight wind speed and direction, aerosol backscatter, depolarization ratio	5 min / 3 m res. up to 10 km range (backscatter) or ~3 km (Doppler velocity)	Doppler velocity < 0.3 m/s Average vertical wind-profile bias to radiosonde: 0.27 m/s
FM-120 Fog Monitor Device (FMD)	Droplet Measurement Technologies (DMT)	20171129 - 20220601	Single-particle forward light scattering	Surface particle diameter, number concentration, LWC, Eff. Diam.	< 1 min / surface obs.	Droplets between 2 – 50 μm
Ceilometer CL51	VAISALA	20171129 - 20220601	Pulsed (6.5 kHz) diode laser Lidar	Cloud intensity, cloud octa and height, aerosol profiles, PBL height	< 1 min / 5 m vert res. up to 15 km a.g.l.	Distance: better than ±5 m
FS11P Visibility Sensor	VAISALA	20171129 - 20220601	Forward-scatter measurement	Visibility, precipitation rate & type, luminescence	< 1 min / surface obs. only	Visibility: ±10% up to 10 km Precipitation sensitivity: 0.05 mm/hr
WXT520 Surface Met. Sensor	VAISALA	20171129 - 20220601	Several sensors & transducers housed in a single unit	2.5 m a.g.l. winds, P, T, RH, and precipitation rate and accumulation	< 1 min / surface obs (a second unit is deployed at 10 m a.g.l.).	Wind speed: ±3% at 10 m/s Wind direction: ± 3° P: ±1hPa (±0.5 hPa 0 to 30 °C) T: ±0.3 °C (at 20 °C) RH: ±3% (0% to 90% RH), ±5% (90% to 100% RH) Accum. precip.: < 5%
Black globe temperature	Campbell Scientific	20171215 - 20220601	Thermistor inside a 6” hollow copper sphere, painted black to measure radiant temp.	Wet bulb globe thermometer (WBGT) index	1 min / surface obs. only	< ± 0.5 °C
Meteorological particle spectrometer (MPS)	DMT	20171215 - 20220601	Optical disdrometer measures droplets from 50 μm to greater than 6.4 mm	2D images of droplets, size distribution, fall velocity, and rain rate	1 min / surface obs. only	Resolution: 25-μm, Range: 50 μm to > 6.4 mm, Concentration range: 0 – 2,000

						particles/cm ³ , Fall speed: < 5%
Parsivel	OTT	20171215 - 20220601	Optical disdrometer that measures liquid/solid particle size and velocity	Particle size, velocity, and precip. rate	1 min / surface obs. only	± 1 size class (0.2 to 2 mm) and ± 0.5 size class (> 2 mm) out of 32 sizes and classes ranging from 0.2 to 25 mm
Pluvio2	OTT	20171215 - 20220601	Precipitation weighing gauge	Precipitation amount and rate	< 1 min / surface obs. Only	± 0.001 mm



750 **Figure 1:** a) Summertime satellite image of Iqaluit, including the CAWS supersite and MSC weather station (red rectangle, image centre-left) located next to the airport runway and south of the smaller T121 site (small red square, top-centre). The inset topographic map shows the locations of Iqaluit (right star) and Whitehorse (left star). b) Same as (a) except for the Whitehorse supersite (small red square, image centre). The red arrow at the top-left in (b) indicates the location of the MSC weather station, just outside of the image's boundary. © Google Earth 2021.

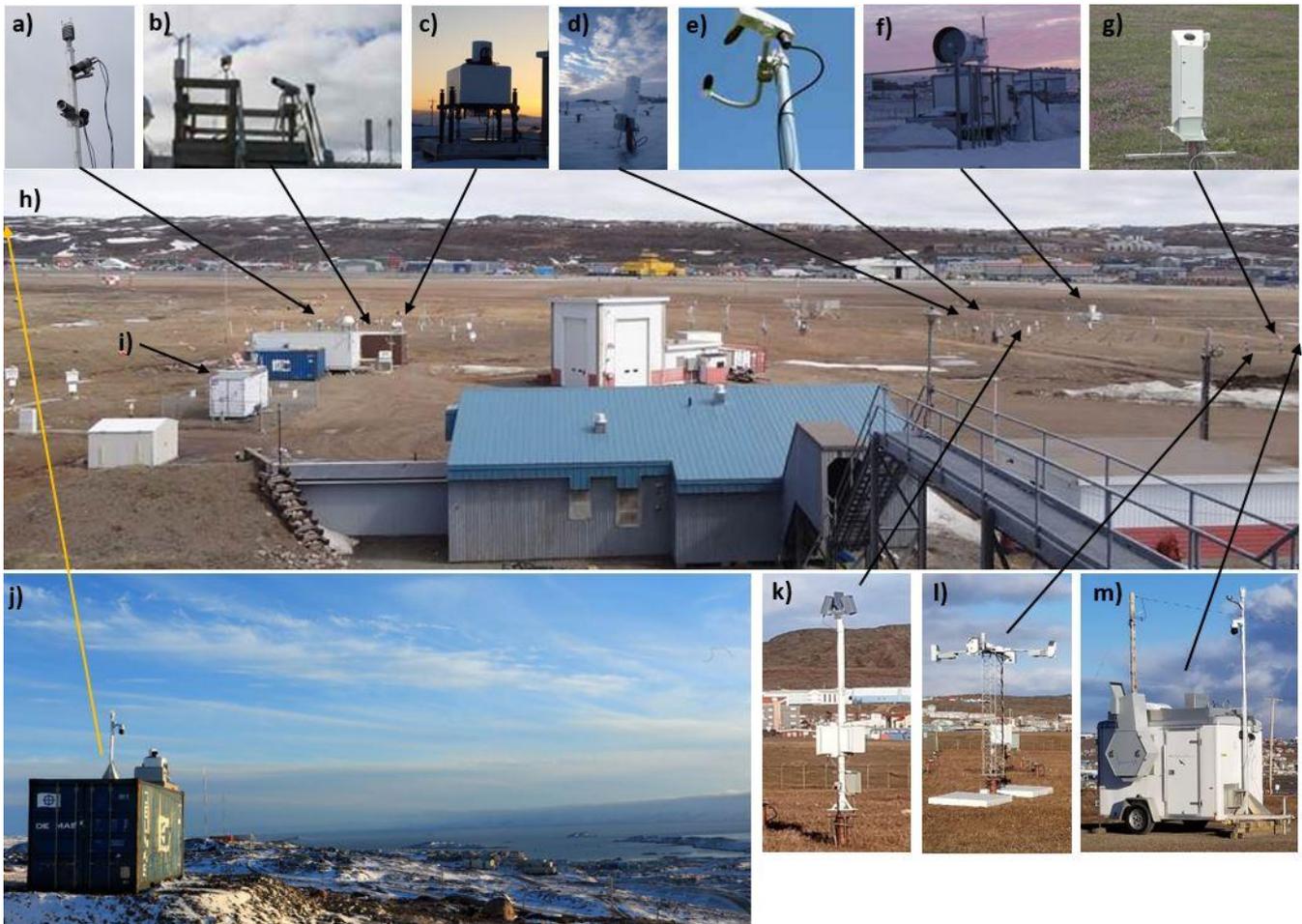


Figure 2: The Iqaluit supersite (h) as viewed in September 2017 from behind the main Weather Station Office (looking East from an elevated platform). The white MSC hydrogen building for radiosonde launches is located in the foreground alongside the standard WMO meteorological observation field. The Iqaluit airport (CYFB) is located in the distance and Frobisher Bay is located to the right (off camera). Most of the CAWS instruments are shown in the inserts: (a) WXT520 (top of the mast) with cameras (below), (b) PIP camera and backlight, (c) Doppler lidar, (d) ceilometer, (e) PWD52, (f) Ka-band radar, (g) DIAL, (h) supersite layout, (i) CAAAL trailer, (j) T121 site layout including a WXT520 and Doppler lidar, (k) POSS, (l) radiation flux sensor suite with soil probe and snow depth sensors, and (m) FIRR (grey instrument) mounted to a trailer. Black arrows indicate each instruments' approximate location; the yellow arrow in (h, j) indicates the approximate location of T121 atop the nearby ridge as seen from the Iqaluit supersite.



Figure 3: Double-fence automated reference (DFAR) configuration at the Iqaluit supersite. Insert: close-up of the single-
770 Alter shielded Pluvio2 precipitation gauge (centre) within inner wooden fence.

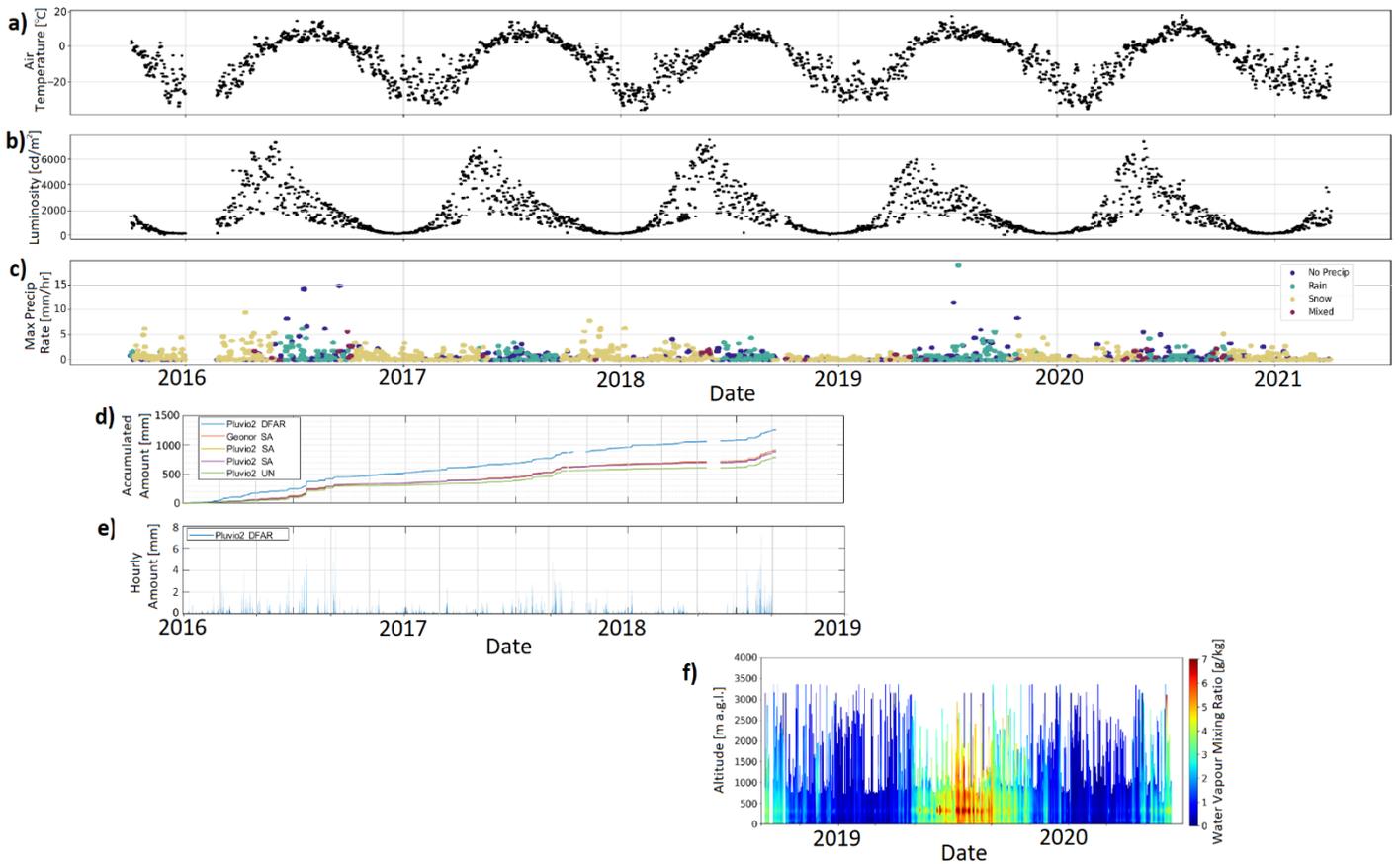


Figure 4: Surface meteorological observations conducted at the Iqaluit supersite from September 29 2015 to April 1 2021. PWD52 daily averaged air temperature (a), luminosity (b), and daily maximum precipitation rate and type (c) are shown for the entire period. Note that mixed precipitation type represents precipitation that is not rain or snow (e.g., freezing rain, sleet, etc.). The C-SPICE's multiple Pluvio2s (DFAR, SA, and UN configurations) and Geonor accumulated precipitation amount (d) and hourly precipitation amount (DFAR Pluvio2 shown only) (e) are shown from January 1 2016 to September 2018, when the C-SPICE project ended. Daily-averaged water vapour mixing ratio profiles observed by the DIAL (f) are shown from the date of its deployment (September 14 2018) until June 20 2020, when the instrument required maintenance.

775

780

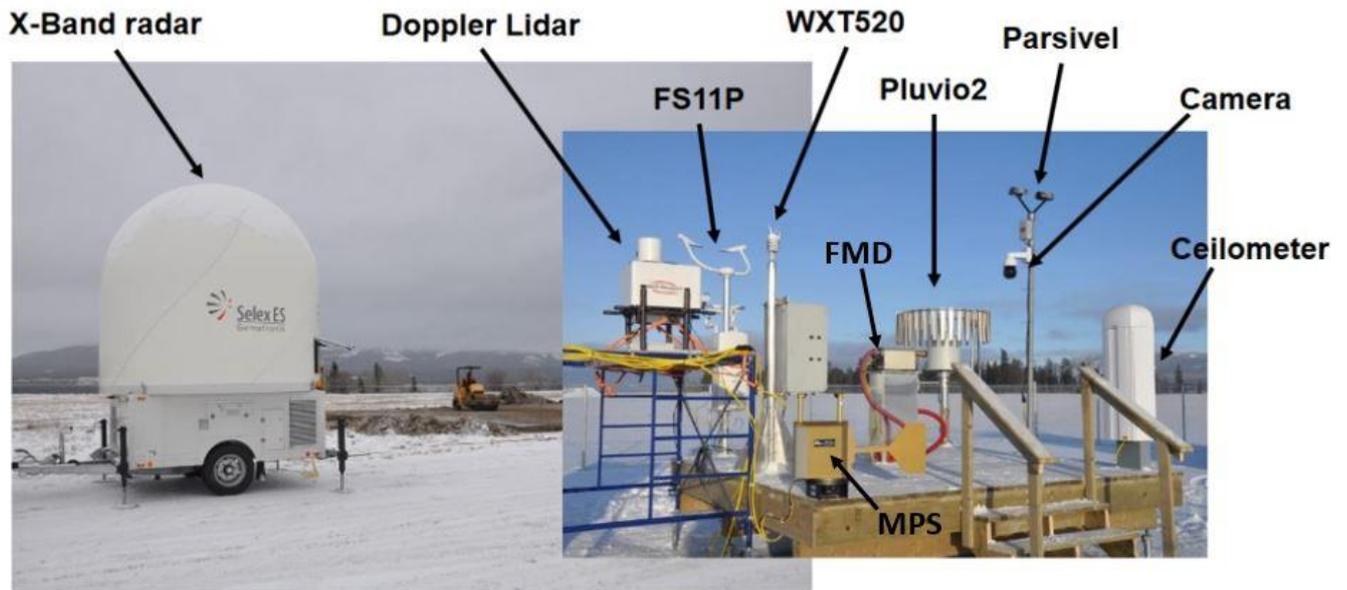


Figure 5: The Whitehorse supersite instrument platform and X-band radar (~10 m away) as viewed on December 12 2017. Images were taken facing NE with the CYXY Whitehorse airport runway behind the photographer. All instruments are labelled.

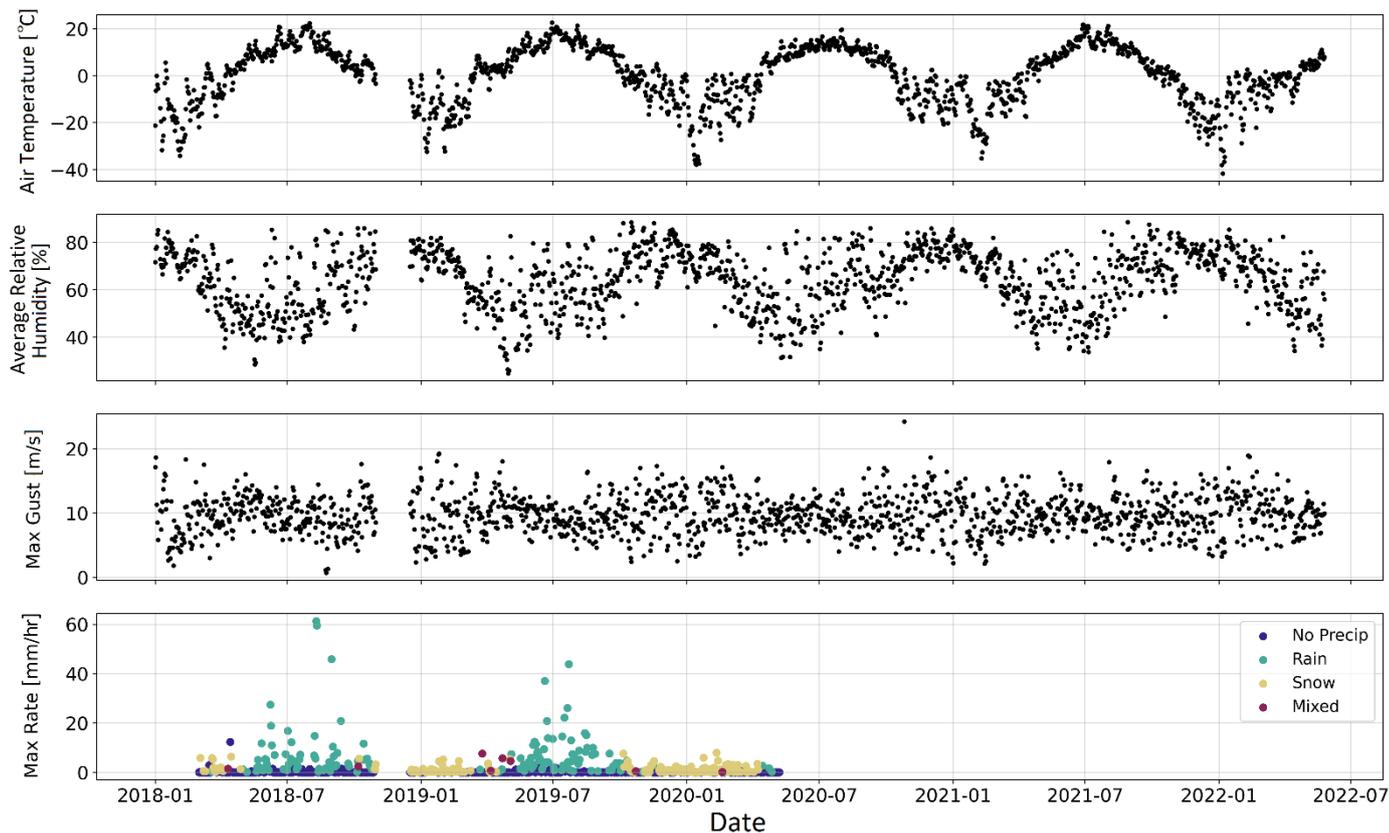


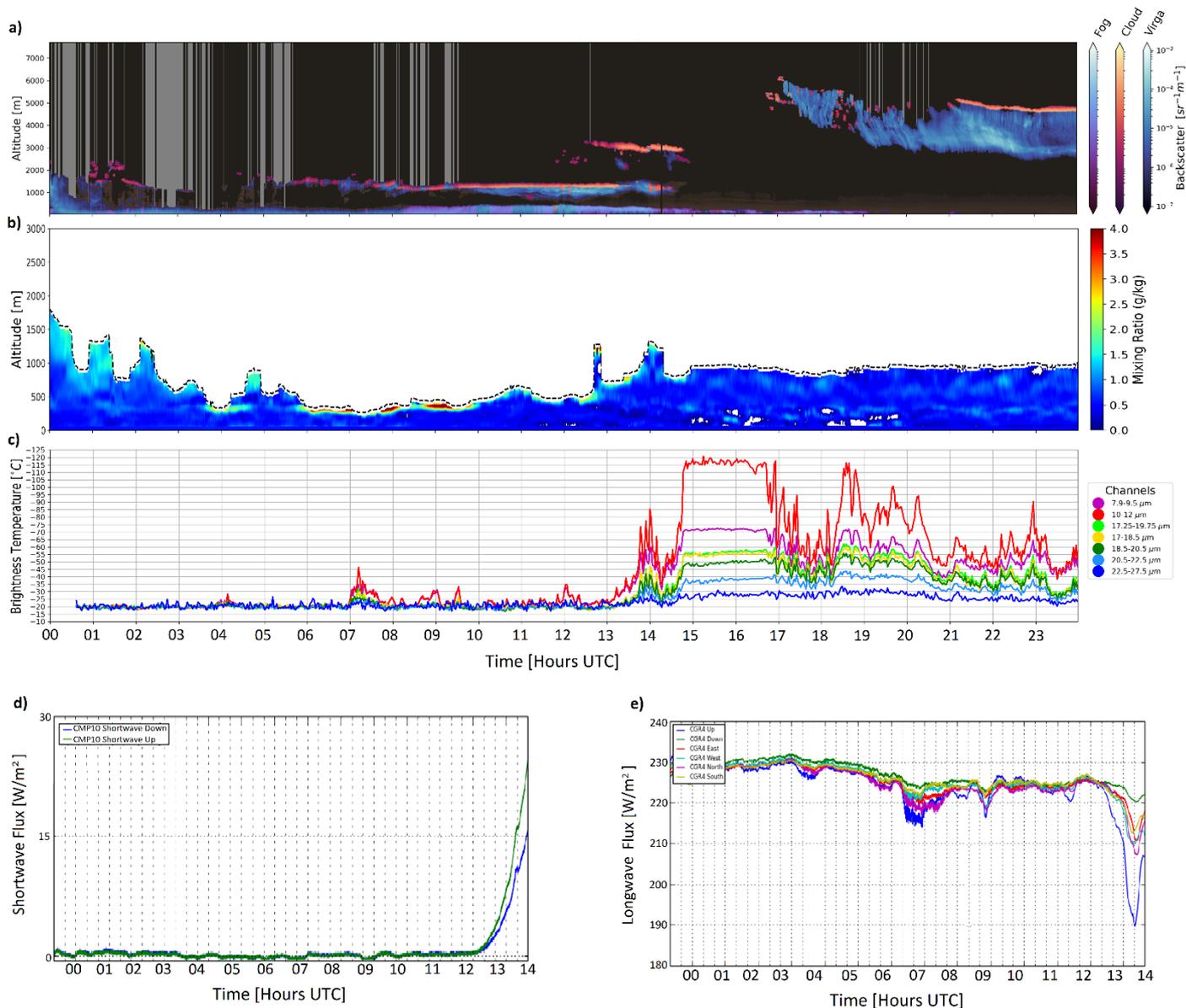
Figure 6: Surface meteorological observations conducted at the Whitehorse supersite during almost the entire study period

(January 1 2018 to June 1 2022). WXT520 daily averaged surface air temperature (a), relative humidity (b), and daily

maximum wind speed (c) are shown. The daily maximum precipitation rate and type (d) were provided by the FS11P, which

790 did not start recording data until February 2018 and experienced a laser failure in June of 2020. The gap of observations in

November 2018 was due to a power outage at the site.



795 **Figure 7:** Multi-instrument observations at the Iqaluit supersite during a blizzard on November 23 2018, including: (a) CL51
 ceilometer aerosol backscatter observations colored by meteorological type classification (clear skies, fog, cloud, and virga),
 (b) DIAL water vapour mixing ratio profiles up to the instruments' effective height (black dashed line), (c) downwelling
 brightness temperature measurements from the FIRR's seven channels, (d) CMP10 shortwave up (green) and downward (blue)
 radiation fluxes, and (e) CGR4 longwave up (blue), downward (green), eastward (red), westward (teal), northward (purple),
 800 and southward (yellow) radiation fluxes. Note (d) and (e) are only shown up to 14:00 UTC due to an instrument malfunction,
 whereas (a-c) are shown for the entire day.

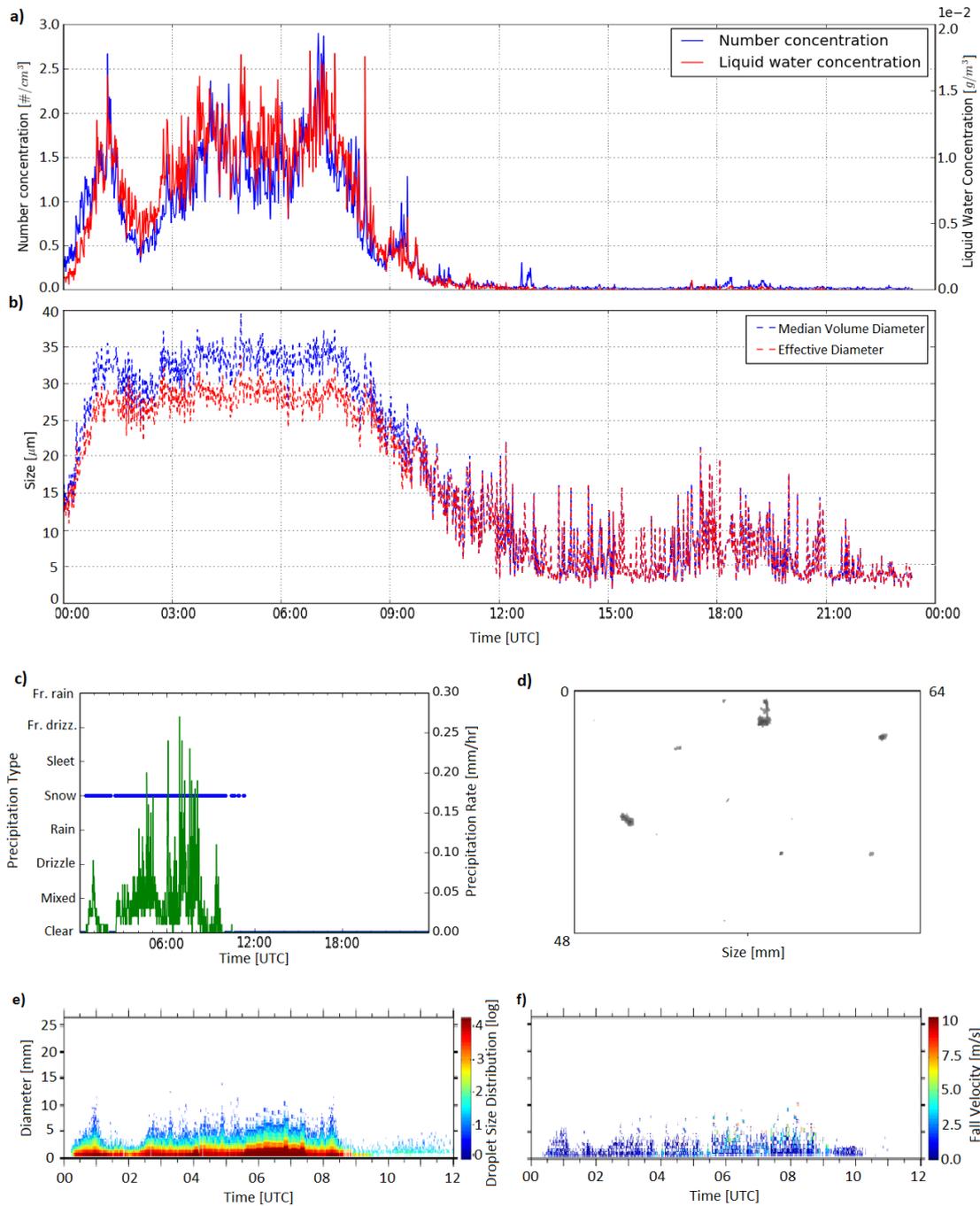
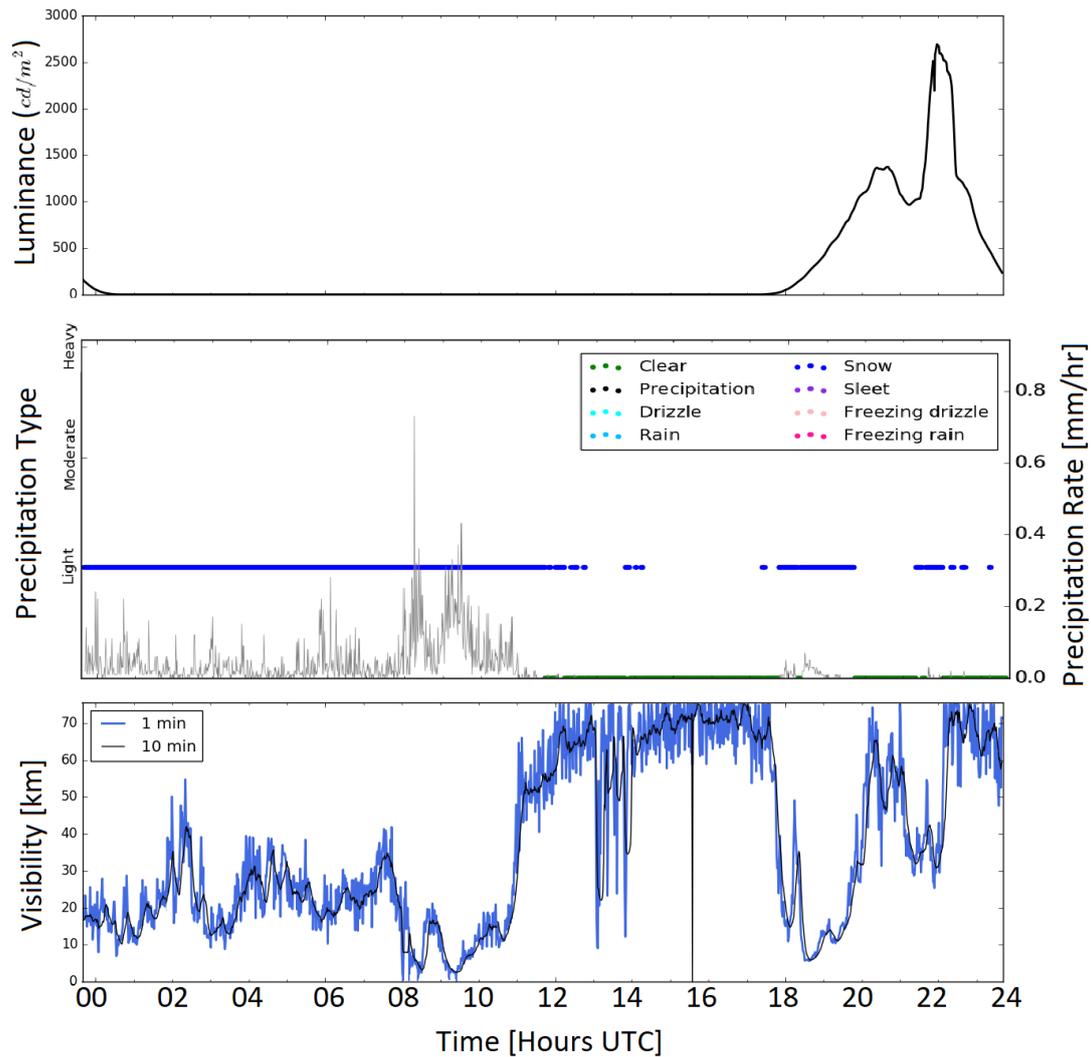
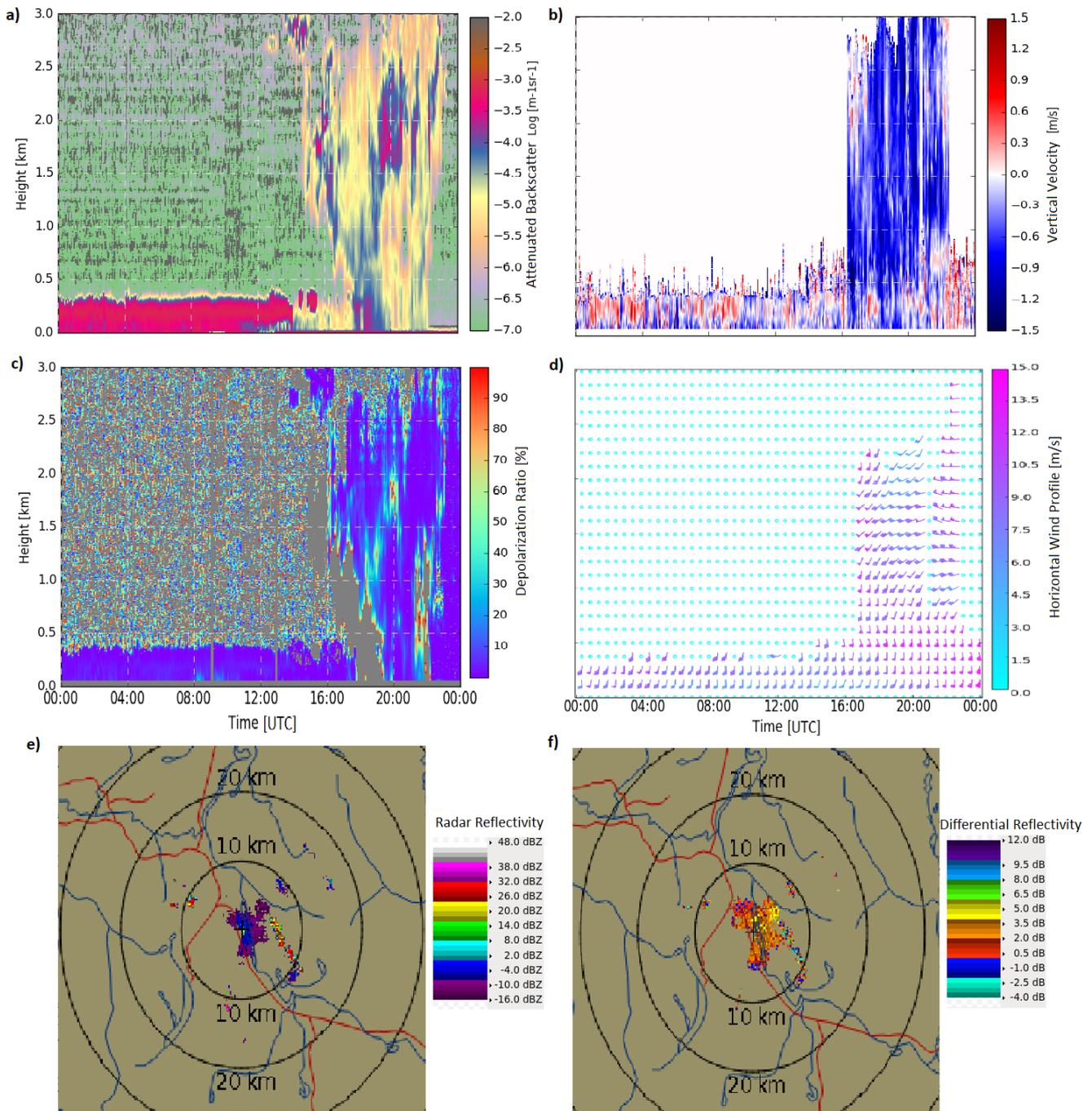


Figure 8: In-situ observations of microphysical properties of precipitation during the November 23 2018 blizzard at Iqaluit:
 805 (a) FMD number (blue) and liquid water (red) concentration, (b) FMD median volume diameter (blue) and effective diameter (red) particle size, (c) PWD52 precipitation rate (green) and type (blue), and PIP observations of (d) snowflakes (photo) taken at 08:37 UTC, (e) DSD, and (f) particle fall velocity during the blizzard.



810 **Figure 9:** FS11P observations at Whitehorse during the blizzard on December 16 2019: (a) luminance, (b) precipitation rate (grey) based on pre-defined intensity thresholds (mm/hr) and type (colour-coded), and (c) Averaged 1-min (blue) and 10-min (black) surface visibility during and after the blizzard.



815 **Figure 10:** Doppler lidar and X-band radar observations during the HIW event at Whitehorse on December 16 2019. Doppler lidar vertical profiles of (a) attenuated backscatter, (b) vertical velocity, (c) depolarization ratio, and (d) horizontal winds are shown. X-band 3.5° elevation PPI scans of radar reflectivity (e) and differential reflectivity (Z_{DR}) (f), provide snapshots of the blizzard in the Whitehorse valley during its peak at 9:45 UTC.