

# Twenty-one years of hydrological data acquisition in the Mediterranean Sea: quality, availability, and research

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10 **Abstract.** Since 2000 and for the following 20 years hydrological data of the Mediterranean Sea, with a particular focus on the western and central Mediterranean sub-basins, have been acquired to study the hydrodynamics at both coastal and open sea scales. Totally 1468 hydrological casts were realized in 29 oceanographic cruises planned due to scientific purposes linked with funding research projects but sometimes driven by sea conditions and type of vessel. After an accurate quality assurance and control, following standard procedures, all hydrological data were included in four  
15 online public open access repositories in SEANOE available from <https://doi.org/10.17882/87567> (Ribotti et al., 2022). Hydrological and dissolved oxygen data are always present in all the datasets whereas pH, fluorescence, turbidity and CDOM are available just for some cruises. Samplings were carried out mainly along transects with some repetitions in the years. The results of two data analyses, staircase systems in the Tyrrhenian Sea and in the Algero-Provencal sub-basin and spreading of the Western Mediterranean Transient, are mentioned.

20 **Keywords:** Mediterranean, CTD, hydrological data, circulation, ocean variability

## 1 Introduction

The Mediterranean basin is an east-west oriented semi-enclosed sea connected with the North Atlantic Ocean, through the narrow and shallow Gibraltar Strait. Its circulation, although characterized by complex patterns, can be simply described as made of three main active overturning circulation cells, a shallow and two deep ones (Tanhua et al. 2013).

25 The Mediterranean Sea is one of the few basins in the world where deep convection and water mass formation take place (Siedler et al., 2001, page 422). Morphology, climatology, and hydrodynamics, with also reduced timescales in the turnover of most of the oceanographic processes of the global ocean, make the Mediterranean Sea a “small and accessible ocean” where changes can be more easily studied (Bethoux et al., 1999).

Within this context, the Italian National Research Council in Oristano and La Spezia (hereafter CNR), organized over  
30 29 scientific cruises in 21 years in the Mediterranean, mostly in the western Mediterranean and the Sicily Channel, with a particular focus on the acquisition of vertical profiles of temperature, salinity, dissolved oxygen, fluorescence and, more recently, pH and turbidity. The cruises were always planned on funding projects’ objectives but then adding further activities linked to other research interests and/or scientific collaborations. Therefore, for each cruise and dataset there is a duality between ‘fit for purpose’, i.e. the projects that prompted us to organize the campaigns, and ‘fit for use’,  
35 i.e. the possibility of reusing the data for purposes other than the original aims, i.e. data also acquired in areas outside projects’ interest. This duality was also behind the sampling plan that could be on a regular or quite regular grid (cruises

2000-2004 and 2014), generally suitable to calibrate and validate ocean numerical models, and/or along transects (after 2004) for studies on general or local circulation and, if with some repetition in the years, for studies on hydrological variability also (for example Gibraltar-Sardinia, Balears-Sardinia, Sardinia-Tunisia) and finally, sometimes linked with the maintenance of the instruments installed on deep moored chains like in Tuscany-Corsica and Sicily-Tunisia. Some cruises or parts of them also show any apparent regular spatial resolution like during a cruise in 2012 with the transect Balears-Sardinia and apparently scattered casts in south Tyrrhenian-Sardinia Channel due to different scientific purposes.

Regardless by any sampling plan, the data acquired over the 2000-2020 period allowed us to characterize the different water masses (Ribotti et al., 2004; Puillat et al., 2006; Santinelli et al., 2008; Belgacem et al., 2020) and the ocean circulation mainly, but not only, in the seas around the Sardinia Island (Sorgente et al., 2016; Pessini et al., 2018, 2020), and to monitor changes with time along repeated transects (Schroeder et al., 2008) or single casts (Durante et al., 2020). But also, to calibrate and validate ocean numerical models at different spatial scales, from open ocean to coastal (Sorgente et al. 2003), also for issues of marine environmental management (oil spill combat; Gonnelli et al., 2016; Sorgente et al. 2016).

The paper is organized as follow: in paragraph 2 (methods), cruises, vessels, and data are described, as well as the quality assurance and check procedures applied to the datasets, subsequently two examples of their use in paragraph 3: the results obtained from data analyses on the diffusion of the Western Mediterranean Transient and a still open issue on thermohaline staircases in the Western Mediterranean are described.

## 2. Methods

Over 29 oceanographic cruises were realized in the Mediterranean basin between May 2000 and October 2020 (table 1 and figure 1) by the CNR in Oristano jointly with the CNR in Spezia, except for the years 2016 and 2019 when no cruises were carried out. The strict collaboration, since 2000 till 2018, with the CNR institute in La Spezia, with a much longer experience in ocean cruises and studies, has brought to the choice of some sampling plans like the transects in the Sicily and Corsica channels where they moored deep oceanographic chains for tens of years.

Totally, 1468 hydrological vertical profiles were acquired in 21 years in the Mediterranean Sea, covering areas spanning from the Gibraltar Strait in the western subbasin to south of Crete in the eastern part of the Mediterranean Sea. In addition to projects' objectives and further research interests, two more aspects have influenced, even heavily in the last four years, data acquisition: sea conditions and type of vessel (described in the next subsection). Until 2017 (table 1), the availability of large research CNR vessels like Urania and Minerva Uno, allowed working throughout the Mediterranean and in "not optimal" sea conditions. Whereas, since 2018 the available research vessels were not suitable to work in rough sea conditions in the open sea.

### 2.1 Research vessels and projects

Different research vessels were used during the several measuring campaigns, each one with characteristics and facilities that influenced both sampling plans and data acquisitions. Between May 2000 and December 2017, the CNR used two multidisciplinary research vessels designed in Italy that permitted to work all over the Mediterranean and in all the seasons: the 61.30 meters long R/V Urania (2000-2014) and the 47,66 meters long R/V Minerva Uno (2015-2017). During these years the performed cruises allowed to acquire data to study the local (e.g. the Sardinia Sea till 2004) and the general circulation all over the Mediterranean, to maintain deep ocean bottom-moored infrastructures in

75 the Corsica, Sardinia and Sicily channels and to acquire data to calibrate and validate numerical models at different spatial scales. In 2016 no cruise was planned.

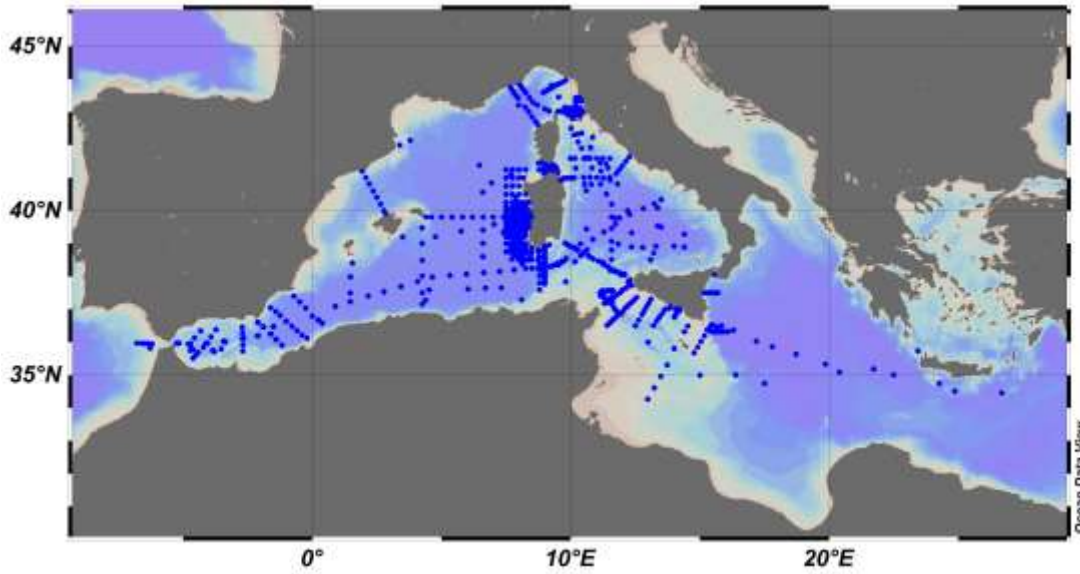


Figure 1. All the CTD casts acquired during the 29 cruises in the Mediterranean between 2000 and 2020.

Cruise name	Start cruise	End cruise	# casts	Vessel	Areas	Ocean parameters
Medgoos1	28/05/2000	02/06/2000	38	Urania	S Sardinia Sea & Channel	P, C, T, O, F
Medgoos2	23/03/2001	03/04/2001	67	Urania	Sardinia Sea & Channel	P, C, T, O, F
Medgoos3	10/09/2001	20/09/2001	41	Urania	S Sardinia Sea & Channel	P, C, T, O, F
Medgoos4	04/05/2002	23/05/2002	68	Urania	Sardinia Sea & Channel	P, C, T, O, F
Medgoos5	31/10/2002	18/11/2002	53	Urania	S Sardinia Channel, Sicily Strait	P, C, T, O, F
Medgoos6	28/03/2003	17/04/2003	92	Urania	S Sardinia Sea & Channel, Sicily Strait, Tyrrhenian Sea	P, C, T, O, F
Medgoos7	07/01/2004	26/01/2004	87	Urania	Sardinia Sea, Sicily Strait, Tyrrhenian Sea	P, C, T, O, F
Medgoos8	08/05/2004	21/05/2004	67	Urania	Tyrrhenian Sea, Corsica Channel	P, C, T, O, F
Medgoos9	06/10/2004	23/10/2004	71	Urania	Gibraltar Strait, Alboran Sea, Algerian basin, Sardinia Channel, Tyrrhenian Sea, Sicily Strait	P, C, T, O, F
Medgoos10	28/05/2005	10/06/2005	42	Urania	Alboran Sea, Algerian basin, Sardinia Channel, Tyrrhenian Sea	P, C, T, O, F
Medgoos11	19/11/2005	01/12/2005	34	Urania	Tyrrhenian Sea, Sicily Strait	P, C, T, O, F
Medgoos12	02/04/2006	17/04/2006	26	Urania	Sicily Channel, Ionian Sea	P, C, T, O, F
Medgoos13	07/10/2006	28/10/2006	56	Urania	Gibraltar Strait, Alboran Sea, Algerian basin, Sardinia & Corsica channels	P, C, T, O, F

MedCO07	05/10/2007	29/10/2007	84	Urania	E Ionian, Sicily Strait, Sardinia Channel, Tyrrhenian Sea, Ligurian Sea, Catalan Sea, Algerian basin	P, C, T, O, F
MedCO08	03/11/2008	24/11/2008	63	Urania	Gibraltar Strait, Alboran Sea, Algerian basin, Sardinia & Corsica channels	P, C, T, O, F
Sicily09	30/10/2009	23/11/2009	109	Urania	channels of Sardinia, Sicily, Corsica, Bonifacio Strait	P, C, T, O, F
Bonifacio2010-Cor	12/03/2010	22/03/2020	28	Maria Grazia	Bonifacio Strait, Corsica Channel	P, C, T, O
Bonifacio2010-Sic	23/11/2010	09/12/2010	22	Urania	Sardinia Channel, Sicily Strait	P, C, T, O
Bonifacio2011	09/11/2011	23/11/2011	19	Urania	Algero-Provençal basin, Tyrrhenian Sea, Sardinia Channel	P, C, T, O, F, Tu
Ichnussa2012	11/01/2012	27/01/2012	36	Urania	Algerian basin, Sardinia Channel, Tyrrhenian Sea	P, C, T, O, F
Ichnussa2013	15/10/2013	29/10/2013	44	Urania	E Ionian, Sicily Strait, Sardinia & Corsica channels, Tyrrhenian Sea, Algerian basin	P, C, T, O, F
Ichnussa2014	13/11/2014	01/12/2014	68	Urania	Tyrrhenian Sea, Sardinia & Corsica channels, Ligurian Sea, Algero-Provençal basin	P, C, T, O, F
SeriousGame2014	17/05/2014	18/05/2014	45	CP 406 A. Scialoja	N Tyrrhenian Sea	P, C, T, O, F, Tu
	21/05/2014	21/05/2014	45			P, C, T, O, F, Tu, CD
Ichnussa2015	25/11/2015	14/12/2015	66	Minerva Uno	Tyrrhenian Sea, Sicily Channel	P, C, T, O, F
Ichnussa2017	23/10/2017	09/11/2017	37	Minerva Uno	channels of Sardinia, Corsica and Sicily, Sicily Strait	P, C, T, O, F
Piattaforme2018	18/05/2018	03/06/2018	29	G. Dallaporta	Sicily Channel, Sicily Strait	P, C, T, O, F, pH, Tu
Ichnussa2018	20/09/2018	03/10/2018	10	G. Dallaporta	Sicily Strait	P, C, T, O, F
Ichnussa2020	09/10/2020	10/10/2020	21	G. Dallaporta	N-NE Sardinia	P, C, T, O, F, Tu

80 **Table 1. List of cruises carried out from 2000 to 2020. Areas with at least three casts are listed. Ocean parameters are listed with acronyms, so P stands for pressure, C for conductivity, T for temperature, O for dissolved oxygen, F for fluorescence, Tu for turbidity, CD for CDOM and pH is just pH.**

In March 2010, the 42.35 m long R/V Maria Grazia was used just for the cruise Bonifacio2010-Cor for a project in the Bonifacio Strait area where an oceanographic and oil spill prediction system based on numerical models was realized  
85 for the local Coast Guard.

During the years 2018-2020, the 35.7 m long R/V Gianfranco Dallaporta, a coastal vessel particularly suitable for activities related to scientific fishing and marine biology, was used in research projects mainly focused on coastal measurements and launch/rescue of Lagrangian surface drifters in south Sicily and Sardinia. In 2019 the 8-days long cruise Ichnussa2019 was cancelled due to bad weather conditions.

90 In May 2014, the 21.5 meters long patrol boat CP406 Antonio Scialoja, a unit of the Italian Coast Guard, was used during a set of two daily surveys named SeriousGame2014-1 and -2 performed in the Tuscan Archipelago, northern Tyrrhenian. Data were acquired to test one of the limited ocean forecast models for the oil spill fate (Gonnelli et al., 2016; Sorgente et al., 2016), part of the wider Mediterranean decision support system (Zodiatis et al., 2016).

## 2.2 Sensors and data

95 For all the adopted vessels, the hydrological instruments were mounted on a General Oceanic rosette system with 24x12-l. Niskin bottles for water column samples collection (2x10-l. during the cruise Bonifacio2010-Cor). This system was equipped with a SBE911 plus CTD probe (Sea-Bird Inc.), sampling the hydrological parameters along the vertical at 24 Hz and with a lowering speed of 1 m/s. With the CTD, data of water conductivity was measured by a SBE-4 sensor, with a resolution of  $3 \times 10^{-4}$  S/m; the water temperature by means of a SBE-3/F thermometer, with resolution of 0.00015 °C/bit at -1 °C or 0.00018 °C/bit at 31 °C; the dissolved oxygen (DO) by means of a SBE-13, with a resolution of 4.3 µM, from Medgoos1 to 4, then using a SBE-43 polarographic membrane sensor with a range of 120% of surface saturation and an accuracy of  $\pm 2\%$  of saturation; the Chlorophyll-a fluorescence, reported as Relative Fluorescence Unit (Chl-a), by means of a Sea Tech Inc. fluorometer with energy emitted by the flash lamp of 0.25 J for flash, with a temperature range between 0 ° - 25 °C and a resolution of 0.15 g/l.

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105 Since the cruise Medgoos12, in April 2006, the Chelsea Aqua 3 fluorometer has been used, whereas, due to sensor malfunction, no fluorescence data were collected during the Medgoos11 and the two cruises Bonifacio2010-Sic and -Cor. Turbidity was acquired just during Bonifacio2011 by a Seapoint, and by a WET Labs ECO-NTU during Piattaforme2018 and Ichnussa2020. The pH/Redox sensor (SBE27) was available just during the cruise Piattaforme2018 but calibrated one year before. During cruises from Medgoos5 to 7 temperature data were checked onboard at defined depths against inverted thermometers, model RTM 4002 by Sensoren Instrumente Systeme GmbH (SiS), installed in correspondence of the Niskin bottles number 1, 3, 5, 7 of the rosette sampler.

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115 Since the cruise Medgoos2 in March-April 2001 (see table 1), redundant or secondary sensors were always used for a data quality assessment (as defined in Bushnell et al., 2019) for measurements of temperature, salinity, and dissolved oxygen (this from Medgoos5 in October 2022; table 1). These secondary sensors were used to evaluate the stability of primaries ones on board and during the following visual data quality check.

Pre-cruise and post-cruise calibrations of first and redundant sensors of temperature, salinity and dissolved oxygen were performed by CNR technicians at the NATO-CMRE Center in La Spezia (Italy). These two calibrations permitted to obtain a slope correction, used in the configuration file of the SBE Seasoft™ suite of programs, improving the data quality. The post-cruise calibration was not performed on data acquired during the cruise Piattaforme2018. After their acquisition, data were pre- processed by the SBE Data Processing™ software with the updated configuration file, to correct coarse errors. Data from the redundant sensor were used instead of the primary in case of malfunction (Ribotti et al., 2020).

120  
125 But, despite their double calibrations, during a cruise all sensors can drift then reducing their data quality. For conductivity and dissolved oxygen sensors, their stability was checked through on-board comparisons with data from water samples. On board the R/V Urania (2000-2014) and R/V Minerva Uno (2015-2017), salinity was checked against the onboard analyses of water, sampled at pre-defined depths, by a Guildline™ 8400B Autosal Laboratory Salinometer, while on the other vessels all samples were analysed once back at CNR (Ribotti et al., 2020). The same for dissolved oxygen measurements that were checked onboard against Winkler titration method (Winkler, 1888) and analysed through the programme TIAMO 2.0™ by Metrohm.

130 Finally, Chl-a, pH and turbidity data were not calibrated. The pressure sensor is usually stable and so it is not calibrated but, in case of malfunction, it was sent to SBE Inc. in USA to be revised.

During all cruises also dissolved inorganic nutrients were acquired from water samples and analyses were carried out once back in the laboratory. Data, from Medgoos8 in May 2004 to Ichnussa2017 in October-November 2017, are available online (Belgacem et al., 2019) and described, in their quality check procedures, in Belgacem et al. (2020).

135 Furthermore, in the framework of the MEDESS4MS project (see info on <https://keep.eu/>), on 17 and 21<sup>st</sup> May 2014 two daily cruises, SeriousGame2014-1 and -2, were jointly held by the CNR in Oristano, the CNR in Ancona and the Italian Coast Guard. The two surveys covered, with 45 casts each, about 880 km<sup>2</sup> of shelf on a 5 x 5 km spatial grid in latitude and longitude of an area located north of the Elba Island in the Tuscany Archipelago, northern Tyrrhenian Sea. The CTD probe was a Seabird SBE19 plus V2 equipped with sensors of pressure, temperature, conductivity, SBE 63 for  
140 dissolved oxygen, a WET-Lab FLNTU for turbidity (NTU) and Chlorophyll (ug/l) and, just during the second cruise, a Turner Cyclops sensor for the Chromophoric Dissolved Organic Matter (CDOM in ppb QS - Quinine Sulfate). The probe was calibrated on January 29<sup>th</sup>, 2013 at SBE Inc. in USA, but not before and after the two cruises, so data were just quality checked and processed by the SBE Data Processing<sup>TM</sup> software to correct the coarse errors.

All CTD data listed in table 1 are available in the SEANOE data repository (Ribotti et al., 2019a, b, c, and Ribotti et al.,  
145 2022) where each dataset is identified per cruise and provided as Ocean Data View (ODV, Schlitzer, 2022) ASCII file format with missing values set to  $-1.e^{10}$ .

Methodologies before, during and after any acquisition, instruments and personnel changed a little during the 21 years of CTD acquisitions, apart from during the two coastal cruises SeriousGame2014-1 and -2 in 2014, due to different adopted instruments. This makes the data in the datasets highly coherent and comparable.

### 150 3. Data analyses

In 21 years, projects and scientific interests led to the choice of areas and sampling plans that could change in relation to vessel characteristics and sea conditions. Nevertheless, repetitions of casts were realized, like in the Sardinia Sea where the same sampling strategy was repeated for 7 cruises in 5 years or the sampling transects between Sardinia and Balears, across the Sicily Strait and the Corsica Channel and along the western Mediterranean, (see in table 1) mainly  
155 for oceanographic studies like circulation, hydrology, and their variabilities. In the following subsections, a process and an oceanographic event, observed from the data described in this paper, are mentioned to suggest further scientific research interests and opportunity for using these data.

In the first subsection we show step structures observed during the first four Medgoos cruises between 2000 and 2002 in the Sardinia Sea, but also during other cruises in the Western Mediterranean. In the following subsection we describe  
160 studies on the Western Mediterranean Transient, that was detected in the datasets acquired between 2005 and 2017.

#### 3.1 Step structures in the deep Sardinia Sea

The first seven cruises named Medgoos have been organized covering the Sardinia Sea, partially or totally. Casts were realized on a regular 25 Km grid from the shelf to the abyssal plain, some hundreds of kilometres far from the coast. The integration of numerical models' solutions and satellite images permitted the study of the circulation in the Sardinia  
165 Sea (Ribotti et al., 2004; Puillat et al., 2006). It is mainly characterized by a fresher modified AW in the upper 200 m, a warmer and saltier Levantine Intermediate Water (LIW) below and till 800 m depth moving northward along the Sardinian slope, and the Western Mediterranean Deep Water (WMDW) between 800 m and the abyssal plain. Due to

baroclinic instabilities of the Algerian Current, surface and intermediate waters are dynamically affected by the presence of deep anticyclonic Algerian Eddies (AEs) along the Sardinian coast, involving surface and intermediate waters till 1500 m of depth (Puillat et al., 2006) and reaching a latitude around 40°N (Puillat et al., 2002; Pessini et al., 2020). But also of the Sardinian Eddies, formed by the interaction of the westward flux of LIW and Tyrrhenian Deep Water (TDW), coming from the Tyrrhenian Sea, with the Algerian Gyre, at the south-western Sardinian corner, influence the circulation of intermediate and deep circulation (Testor and Gascard, 2005; Testor et al., 2005). Cruises Medgoos2, 4 and 7 (on March 2001, May 2002, and January 2004, respectively; see table 1) covered the whole Sardinia Sea permitting to investigate some hydrological profiles of March 2001 and repeat some casts hundreds of kilometres far from Sardinia. AEs were observed during the first two cruises in the southern-central part of the Sardinia Sea (Puillat et al., 2006), while in January 2004 the area was eddy-free. The profiles from these three cruises showed step structures at latitudes between 40.5° and 41.5°N, offshore north-western Sardinia at depths between 500 m and 1400 m. Following the method proposed by Durante et al. (2019) to investigate the staircases systems in the central Tyrrhenian Sea, the analysis of data collected during Medgoos cruises revealed always a 5 to 9 steps structure with interfaces of 50-110 m thick located between intermediate and deep waters as depicted by the profiles of absolute salinity and conservative temperature reported in figure 2.

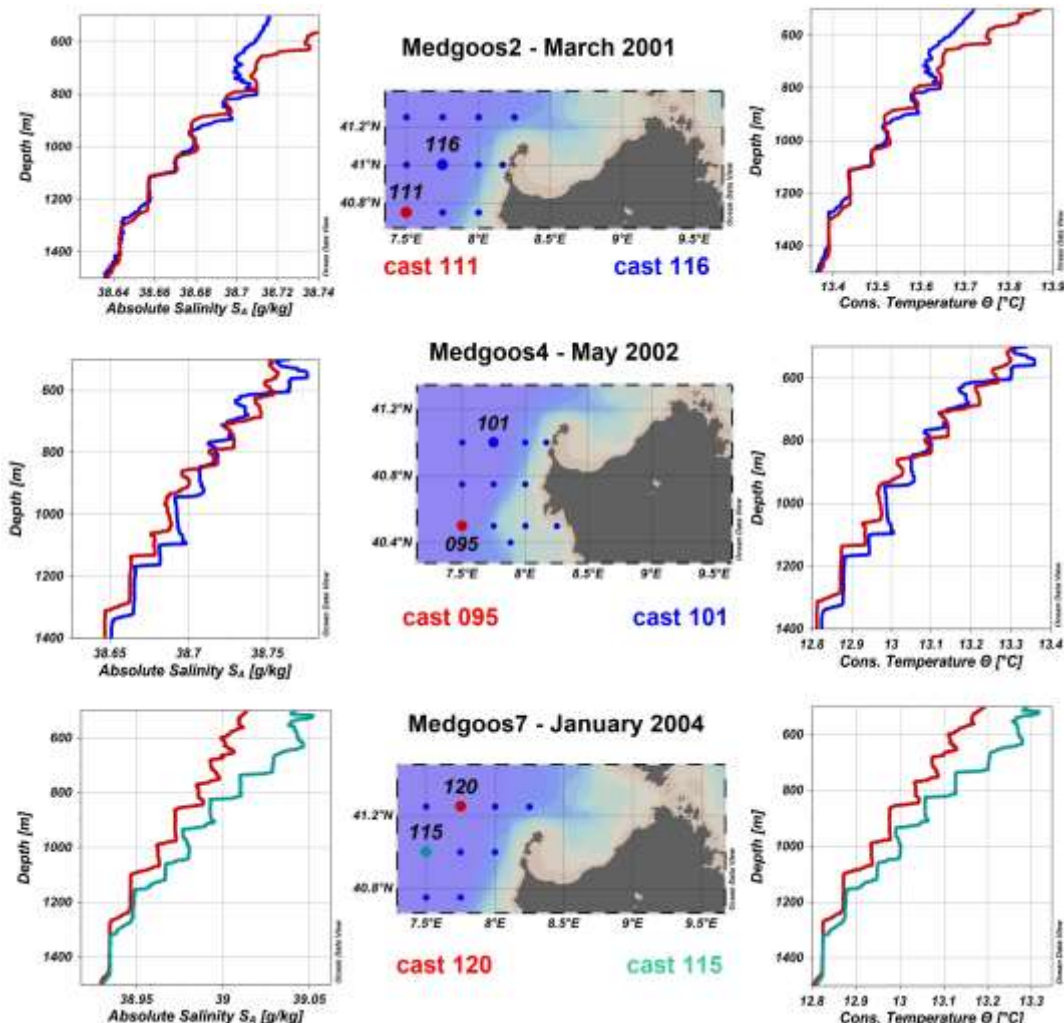


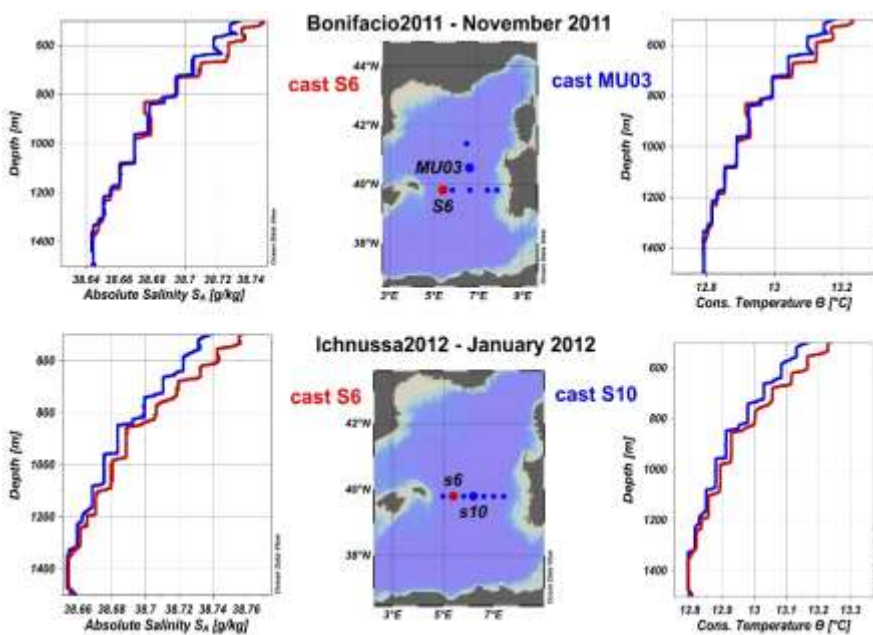
Figure 2. Some of the steps observed offshore north-west Sardinia between 500 and 1500 m through the profiles of absolute salinity and conservative temperature at the interface of intermediate-deep waters in the period March 2001 - January 2004.

During the first cruise conducted in March 2001 the staircase structure was less evident than in 2002 and 2004. Each step is defined by a layer with values of the stability ration ( $R_\rho$ ) and of the Turner angle ( $Tu$ ) finger regime, both defined



in Durante et al. (2019), comprised between 1 and 1.7 ( $1 < R_p < 1.7$ ) and between  $45^\circ$  and  $90^\circ$  ( $45^\circ < Tu < 90^\circ$ ), respectively. Following Durante et al. (2019), these staircases systems can be defined with rough profiles, then representing instabilities in the convective layer, and characterized by a reduced number of thick steps alternated by rough interfaces. The ranges of the differences for temperature and salinity at the interface of each step ranged between  $1 \cdot 10^{-2}$  and  $1.5 \cdot 10^{-1} \text{ }^\circ\text{C}$  and between  $1 \cdot 10^{-2}$  and  $9 \cdot 10^{-3}$  PSU, respectively, with an interface thickness varying between 9 m and 40 m. These values are comparable with those observed in the Tyrrhenian Sea by Zodiatis and Gasparini (1996) with the difference that the central Tyrrhenian Sea is a low energy area in terms of water currents, whereas the northern Sardinia Sea is at the limits of the passage of deep anticyclonic eddies moving anti-clockwise in the Algerian sub-basin and generating intense currents (Pessini et al., 2018, 2020).

Staircases systems were also found in the central Algero-Provencal basin during the Bonifacio2011 (November 2011) and Ichnussa2012 (January 2012) cruises and were well identified from the analysis of the profiles of absolute salinity and conservative temperature profile as reported in figure 3.



200 **Figure 3. Profiles of absolute salinity and conservative temperature of two casts where the steps were observed along the transect between Balears and Sardinia at 500-1500 m depth in the period November 2011 – January 2012.**

These staircases are organized in 7 to 9 steps structures with interfaces 60-110 m thick and  $R_p$  and  $Tu$  comparable with those observed offshore north-west Sardinia. These systems were detected in sampling stations located in the central Provencal sub-basin and in the central part of the transect between Minorca Island in the Balears and Oristano in Sardinia, along latitude at around  $40^\circ$  N. They are similar to those described by Fuda et al. (2000) observed between the 500 and 1400 m depths from CTD and XBT profiles acquired in February 1994 in the open Algerian area.

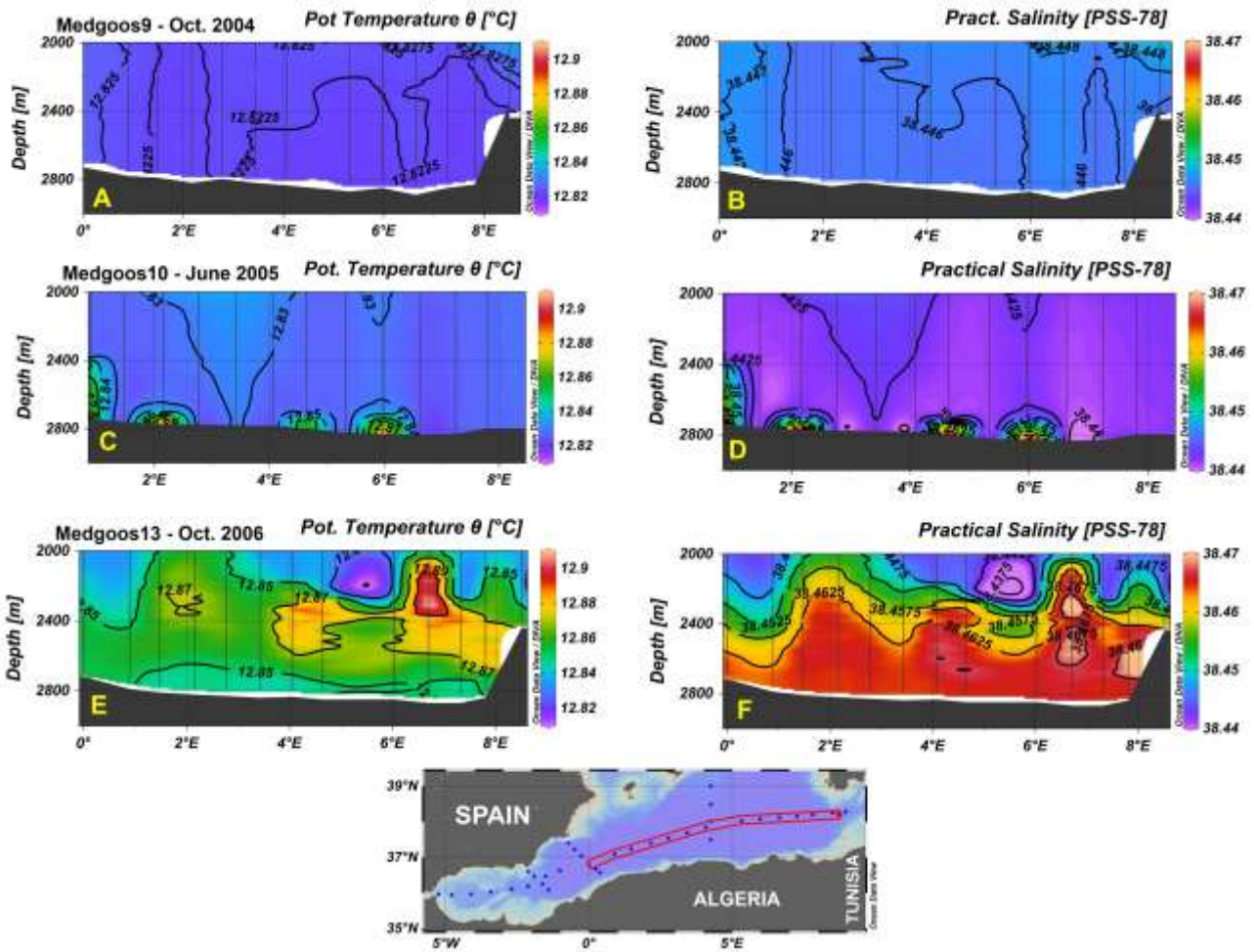
The features of these dynamical structures confirm they are related to salt fingering originated by the presence of intermediate waters, both when close to the continental slopes (NW Sardinia), or when in the middle of the Algero-Provencal area. In both cases, their presence is probably dependent on the absence of eddies that, if deep, can erode the steps' structure, in agreement with what observed in June 2017 by Taillandier et al. (2020) in the Algerian Basin along the  $38^\circ$  N parallel. Finally, during 12 cruises between March 2003 (Medgoos6) and December 2015 (Ichnussa2015), the 3550 m deep cast number 51 in the central Tyrrhenian Sea was repeated to study the staircase system previously mentioned and recently well described by Durante et al. (2019).



215 **3.2 Deep changes in the western Mediterranean sub-basin**

In the Mediterranean the deep circulation is organized in two deep cells, located in the western and eastern sub-basin respectively, and constituted by denser waters than those above whose thermohaline features can differ from between the two sub-basins (Malanotte-Rizzoli et al. 1997; Sparnocchia et al., 1999; Wu et al., 2000; Bethoux et al., 2022; Rixen et al., 2005; Marty and Chiavérini, 2010; Bensi et al., 2013; Ingrosso et al., 2017; Send and Testor, 2017).

220 Between October 2004 (Medgoos9) and 2017 (Ichnussa2017), fifteen cruises covered some deep sub-basins of the western and central Mediterranean, from the Sicily Channel and the Tyrrhenian Sea over the Gibraltar Strait, with hydrological casts (see table 1). These data, apart the shallow cruise Bonifacio2010-Cor and the two SeriousGame in 2014, permitted to follow the spreading of the new Western Mediterranean Deep Waters (nWMDW) in the western Mediterranean sub-basins since its formation in 2005 (Schroeder et al., 2008), revealing over a decade of great  
225 thermohaline variability. During winter 2004/2005, above the north-western Mediterranean area, a little rain with intense and persistent northerly winds brought a massive heat loss in the Gulf of Lions, the highest since 1948, that led to an increase in surface salinity (Font et al., 2007). In the same years, the propagation of the Eastern Mediterranean Transient, from the eastern to the western basin of the Mediterranean, led to maximum values of temperature and salinity in the intermediate layers (Gasparini et al., 2005; Lopez-Jurado et al., 2005; Schroeder et al., 2006; Margirier et al., 2020). The concomitance of these meteorological and oceanographic processes has created the perfect conditions  
230 for the formation of warmer and saltier deep waters than the pre-existing ones (Herrmann et al., 2010; Lopez-Jurado et al., 2005), then occupying the deepest layer close to the bottom and known as Western Mediterranean Transient (WMT; Lopez-Jurado et al., 2005; Schroeder et al., 2008; Briand, 2009; Zunino et al., 2012). After the first event of 2004/05, WMT continued with large volumes of nWMDW formed in winter 2005/06 (Medgoos10&13 in figure 4) while the  
235 following 2006/07 and 2007/08 no deep water production occurred due to a mild “cold” season. But deep convection started again in winter 2008/09 and in the following (i.e. 2009/10, 2010/11 and 2011/12) due to particularly cold winters, like 2011/12, in the Mediterranean Sea. In 2012 and at the bottom of the Gulf of Lions, the nWMDW had a potential temperature of 12.905 °C, a salinity of 38.508 and a maximum density of 29.134 kg m<sup>-3</sup> (Durrieu de Madron et al., 2013). Six years earlier and at the bottom of the Algerian basin (Medgoos13 in 2006), values were lower with  
240 12.845 °C, 38.465 and 29.115 kg m<sup>-3</sup> of potential temperature, salinity and density, respectively, as shown in figures 4E, F and 5C.



245 **Figure 4.** Deep sections (over 2000 m depth) of potential temperature ( $^{\circ}\text{C}$ ) and practical salinity (dimensionless) along a transect at  $37\text{--}37.5^{\circ}\text{N}$  between the eastern entrance to the Alboran Sea ( $0^{\circ}\text{E}$ ) and the western entrance to the Sardinia Channel ( $8.5^{\circ}\text{E}$ ) during the cruises Medgoos9 (A-B; October 2004), before the WMT, Medgoos10 (C-D; May-June 2005) and Medgoos13 (E-F; October 2006) with the filling of the Algerian sub-basin by the WMT.

Every winter, new warmer, saltier, and denser deep waters formed, moving to a continuous increase in heat and salt content in the deep layer, making the deep temperature-salinity ( $\theta\text{S}$ ) diagrams more and more complex (Schroeder et al., 2016). The replacement of the old deep waters near the bottom by new ones, is visible as a creep close to the bottom in 2005, becoming 600 m thick in 2006, almost 1000 m thick in 2008, over 1200 m in 2010, 1400 m in 2013 and  $> 1500$  m in 2015. Since 2010 the signal of the WMT was found also inside the Tyrrhenian Sea, after having overpassed the Sardinian Trough, connecting the Algerian sub-basin with the Tyrrhenian one along the axis of the Sardinia Channel. The WMT changed structure and properties of the deep layers in the Western Mediterranean recognizable by hooks and inversions in characteristic  $\theta\text{S}$  diagrams (in figure 5 in comparison with Medgoos9 data, prior the WMT). This set of CTDs acquired during over a decade (2004–2017) can contribute to the understanding of the temporal and spatial evolution of the thermohaline variability in the western Mediterranean Sea. The thermohaline circulation of the Mediterranean Sea, through the Mediterranean Outflow Water (MOW), has a particular impact on the North Atlantic circulation and the Atlantic Meridional Overturning Circulation (AMOC) due to its high salinity, as rebuilt in palaeoceanographic studies (Rogerson et al., 2012; van Dijk et al., 2018; Ausín et al., 2020; Mesa-Fernández et al., 2022).

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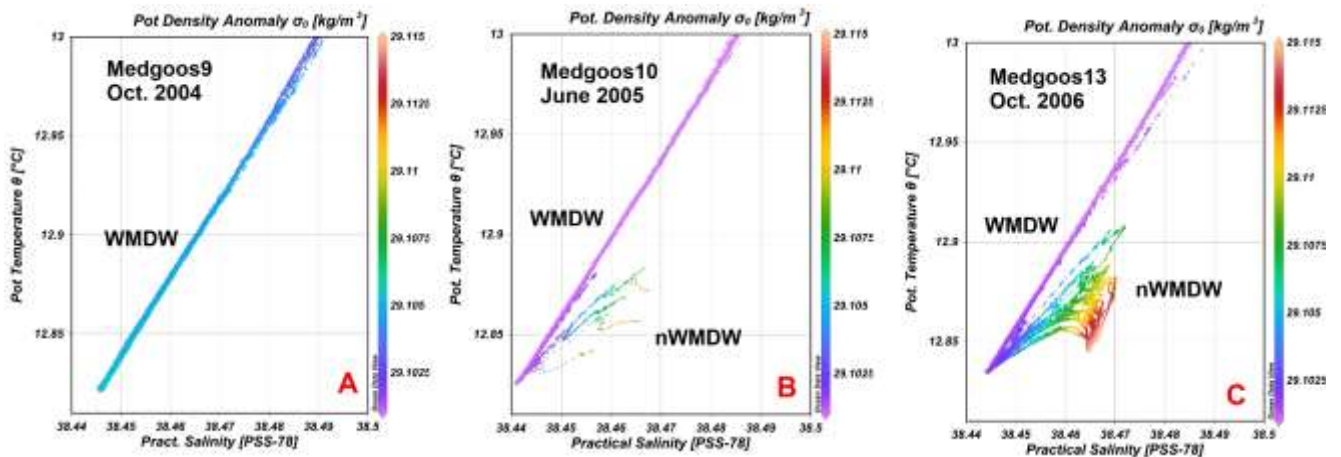


Figure 5. Potential temperature – practical salinity ( $\Theta_S$ ), with potential density anomaly ( $\sigma_\theta$ ; Kg/m<sup>3</sup>) of the casts along the transect of figure 4 during the cruises Medgoos9 (A; October 2004), before the WMT, Medgoos10 (B; May-June 2005) and Medgoos13 (C; October 2006) with the typical hooks and inversions of the new deep waters' characteristics.

#### 265 4. Data availability

The four datasets described in this study are publicly available and free of charge from the SEANOE data repository as Ribotti et al. (2019a, b, c) in <https://doi.org/10.17882/59867>, <https://doi.org/10.17882/70340> and <https://doi.org/10.17882/59777> respectively, and as Ribotti et al. (2022) in <https://doi.org/10.17882/87567>. The presented datasets are composed of CTD data in Ocean Data View (ODV) TXT Spreadsheet files, one for each cruise.

#### 270 5. Discussion and conclusion

Over 1468 hydrological vertical profiles were acquired during 29 oceanographic cruises in 21 years in the Mediterranean Sea, with a particular focus on the western and central Mediterranean sub-basins, between May 2000 and October 2020. Each cruise was mainly led by the projects' objectives funding the research but also by scientific interests at different spatial (coastal or offshore) and temporal (from days to years) scales. Use and reuse of data with time was mainly driven by the following issues: solution of marine environmental problems, studies on ocean circulation and variability, and/or model validation. This is visible in the different spatial resolution of the sampling plans, on a regular grid of casts or along repeated transects, sometimes both used during the same cruise.

During the years and despite of some improvements in sensors quality, sensors and instruments were prepared through the same procedures before and, after their use, the same quality assurance protocols, apart in the two Serious Games cruises in 2014. CTDs data followed all quality check, assessment, standard and best practices, defined at international level (see Bushnell et al., 2019) and through standardized procedures for all sensors (Hood et al., 2010). Pre- and post-calibration, redundant sensors, on-board analyzed water samples has achieved the best accuracy standard for the used instruments. The presence of the same operators on-board in most of the years has then assured a reduced uncertainty of measurements further increasing the quality of the data set.

285 The final results are coherent datasets of CTD data that can be used for all studies, as described in the paper for the WMT or for salt fingering.

## Author contribution

AR led projects funding data acquisition and cruises, finalized data quality procedures described in the paper, realised datasets and the writing of the paper. MB led cruises, finalized data quality procedures described in the paper, realised datasets and collaborated to the paper writing. RS and AC led projects funding data acquisition and collaborated to the data analyses and paper writing. FP and QC collaborated to the data analyses and paper writing.

## Competing interests

The authors declare that they have no conflict of interests.

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