



1 **AnisoVeg: Anisotropy and Nadir-normalized MODIS MAIAC datasets for satellite**
2 **vegetation studies in South America**

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26 **This document includes:** main manuscript, figures and tables.



27 Abstract

28 The AnisoVeg product consists of monthly 1-km composites of anisotropy (ANI) and nadir-
29 normalized (NAD) surface reflectance layers obtained from the Moderate Resolution Imaging
30 Spectroradiometer (MODIS) sensor over the entire South America. The satellite data were pre-
31 processed using the Multi-Angle Implementation Atmospheric Correction (MAIAC). The
32 AnisoVeg product spans 22 years of observations (2000 to 2021) and includes the reflectance of
33 MODIS bands 1 to 8 and two vegetation indices (VIs): Normalized Difference Vegetation Index
34 (NDVI) and Enhanced Vegetation Index (EVI). While the NAD layers reduce the data variability
35 added by bidirectional effects on the reflectance and VI time series, the unique ANI layers allow
36 the use of this multi-angular data variability as a source of information for vegetation studies. The
37 AnisoVeg product has been generated using daily MODIS MAIAC data from both Terra and Aqua
38 satellites, normalized for a fixed solar zenith angle ($SZA = 45^\circ$), modelled for three sensor
39 view directions (nadir, forward, and backward scattering), and aggregated to monthly composites.
40 The anisotropy was calculated by the subtraction of modelled backward and forward scattering
41 surface reflectance. The release of the ANI data for open usage is novel, as well as the NAD data
42 at an advance processing level. We demonstrate the use of such data for vegetation studies using
43 three types of forests in eastern Amazon with distinct gradients of vegetation structure and
44 aboveground biomass (AGB). The gradient of AGB was positively associated with ANI, while
45 NAD values were related to different canopy structural characteristics. This was further illustrated
46 by the strong and significant relationship between EVI_{ANI} and forest height observations from the
47 Global Ecosystem Dynamics Investigation (GEDI) LiDAR sensor considering a simple linear
48 model ($R^2 = 0.55$). Overall, the time series of the AnisoVeg product (NAD and ANI) provide
49 distinct information for various applications aiming at understanding vegetation structure,
50 dynamics, and disturbance patterns. All data, processing codes and results are made publicly
51 available to enable research and the extension of AnisoVeg products for other regions outside the
52 South America. The code can be found at <https://doi.org/10.5281/zenodo.6561351> (Dalagnol and
53 Wagner, 2022), EVI_{ANI} and EVI_{NAD} can be found as assets in the Google Earth Engine (GEE)
54 (described in the data availability section), and the full dataset is available at the open repository
55 <https://doi.org/10.5281/zenodo.3878879> (Dalagnol et al., 2022).

56 **Key-words:** AnisoVeg, South America, vegetation structure, forest monitoring, MODIS.

57

58 1. Introduction

59 The anisotropy is defined by the directional dependence of observations on mechanical or
60 physical properties of surfaces. Because most land covers are not Lambertian (isotropic), the
61 surface reflectance measured by satellite sensors varies with the view zenith angle (VZA), view
62 direction (backward or forward scattering), and solar zenith angle (SZA) (Galvão et al., 2011).
63 This is especially valid for images acquired over vegetated surfaces by large field-of-view (FOV)
64 instruments such as the Moderate Resolution Imaging Spectroradiometer (MODIS) (Bhandari et
65 al., 2011). MODIS has a wide swath scanning $\pm 55^\circ$ from nadir on board the Terra and Aqua
66 satellites. For example, a reflected signal coming from the backward scattering direction of
67 MODIS under a large VZA and close-to-zero relative azimuth angle (RAA) between the satellite
68 and sun (sun behind the platform) is generally higher than that coming from the nadir ($VZA = 0^\circ$)
69 or forward scattering direction (platform facing the sun at $RAA = 180^\circ$). Moreover, the SZA also
70 varies seasonally and across geographical locations, affecting the amounts of shadows in the
71 surfaces observed by satellites (Galvão et al., 2013). Such view-illumination effects are dependent
72 on the land cover types and their magnitude relates to differences in biophysical properties of the
73 vegetation (Foody & Curran, 1994). Therefore, the vegetation anisotropy can be seen
74 antagonistically as sources of noise and biophysical information in the time-series analysis of



75 vegetation indices (VIs) calculated from MODIS. As a source of noise, one may consider that the
76 reflected signal toward the large FOV satellite sensors varies with distinct view-illumination
77 geometries of data acquisition over the same surface. As a source of information, one may
78 highlight that the anisotropy is land-cover type dependent, showing spectral variations that may
79 be associated, for instance, with changes in vegetation structure across different forests.

80 To reduce the bidirectional effects as a source of noise, a nadir-normalized dataset can be created.
81 We can normalize the surface reflectance of the MODIS bands to a specific set of VZA and SZA
82 using the bidirectional reflectance distribution function (BRDF), represented by a model such as
83 the Ross-Thick Li-Sparse (RTLS) (Wanner et al., 1995). To ensure confidence in the data
84 analysis, we can also use the Multi-Angle Implementation Atmospheric Correction (MAIAC) for
85 atmospheric correction. MAIAC is a new generation of cloud screening and atmospheric
86 correction algorithm that uses an adaptive time series analysis and processing of groups of pixels
87 to derive atmospheric aerosol concentration, cloud mask and surface reflectance without typical
88 empirical assumptions (Lyapustin et al., 2011, 2012). It offers substantial improvement over
89 conventional algorithms by mitigating atmospheric interference and advancing the accuracy of
90 surface reflectance over tropical vegetation by a factor of 3 to 10 (Hilker et al., 2012). Due to the
91 improvements in cloud detection, aerosol retrieval and atmospheric correction, the MAIAC
92 algorithm provides from 4 to 25% more high-quality retrievals than the traditional MOD09
93 product, with the largest estimate being observed for tropical regions (Lyapustin et al., 2021).
94 Studies have used MODIS MAIAC observations with nadir-normalized geometry to assess
95 Amazonian forests' structure, functioning, and impacts of environmental and climate change
96 (Hilker et al., 2014; Wagner et al., 2017; Anderson et al., 2018; Dalagnol et al., 2018; Fonseca et
97 al., 2019; Bontempo et al., 2020; Gonçalves et al., 2020; Zhang et al., 2021). For instance, such
98 product provided reliable time series of surface reflectance data that allowed to identify large-
99 scale communities of bamboo species and their dynamics in the southwest Amazon (Dalagnol et
100 al., 2018). Lastly, by improving the cloud screening and minimizing BRDF artifacts in
101 comparison to uncorrected data, the MAIAC greatly contributed to the understanding of the long-
102 standing debate in the Amazon over the possible existence of the green-up phenomenon observed
103 during the dry season of each year or with severe droughts (Morton et al., 2014; Bi et al., 2015;
104 Saleska et al., 2016; Wu et al., 2017). The existence of this phenomenon has implications on the
105 comprehension of the resilience of tropical forests to climate change.

106 To use the bidirectional effects as a source of information, we generate an anisotropy dataset that
107 is dependent on land-cover types and captures the variations of sunlit and shaded canopy
108 components viewed by the sensors (Chen et al., 2003; Gao, 2003). The use of multi-angular
109 information to obtain metrics of anisotropy and extract information on forest structure was
110 suggested two decades ago (Foody & Curran, 1994). The first experiments with such concept
111 were conducted by calculating the ratio between backward and forward scattering data and
112 generating the anisotropy index (ANIX) on studying short-stature grass-type vegetation
113 (Sandmeier et al., 1998). Other indices have been developed and validated afterwards (Schaaf et
114 al., 2002; Lacaze et al., 2002; Chen et al., 2005; Pocerwicz et al., 2007; Moura et al., 2015; Sharma
115 et al., 2021). However, this remains an understudied topic with limited results reported in the
116 literature, especially in tropical regions. For instance, observations from the Multi-angle Imaging
117 Spectroradiometer (MISR)/Terra in the backward and forward scattering directions facilitated the
118 discrimination of savanna physiognomies in Brazil (Liesenberg et al., 2007). MODIS MAIAC
119 data from both directions were also used to calculate an anisotropic VI that explained part of the
120 large-scale photosynthetic activity in the Amazon, where higher photosynthetic activity was
121 associated to higher anisotropy values (Sousa et al., 2017). Moura et al. (2015) employed a more
122 sophisticated approach based on scattering at backward and forward view directions using multi-
123 temporal and multi-angular observations of MAIAC MODIS and BRDF modelling. The resultant
124 metrics of anisotropy were further validated against field and airborne Light Detection And



125 Ranging (LiDAR) observations, showing strong linear relationship with leaf area index (LAI) (R^2
126 = 0.70-0.88), canopy heterogeneity ($R^2 = 0.54$), and photosynthetic activity ($R^2 = 0.73-0.98$)
127 (Moura et al., 2015; Moura et al., 2016; Hilker et al., 2017). Although showing great potential in
128 vegetation studies, the aforementioned anisotropy metrics were never computed over larger areas
129 of the world such as proposed in this study for South America.

130 The objective of this work is to present the AnisoVeg product, and how it can be used for
131 vegetation studies. We use MODIS Collection 6 (C6) MAIAC (Lyapustin et al., 2018) monthly
132 data (2000-2021) generated at 1-km spatial resolution for the entire South America with two
133 different types of layers: (1) nadir-normalized (NAD) data for the surface reflectance of MODIS
134 bands 1 to 8 and two VIs (NDVI and EVI); and (2) anisotropy data (ANI) calculated from the
135 difference between backward and forwarding scattering estimates of bands 1 to 8 and VIs (Moura
136 et al., 2015). The motivations for generating this product extend from developing applications of
137 multi-angle observations for vegetation studies to producing analysis-ready and openly available
138 datasets of anisotropy and nadir metrics for a larger community of users. The paper is organized
139 in several sections to present the processing steps for generating the AnisoVeg products, a brief
140 evaluation of data products over experimental areas, and finally an example of its potential
141 application in vegetation studies.

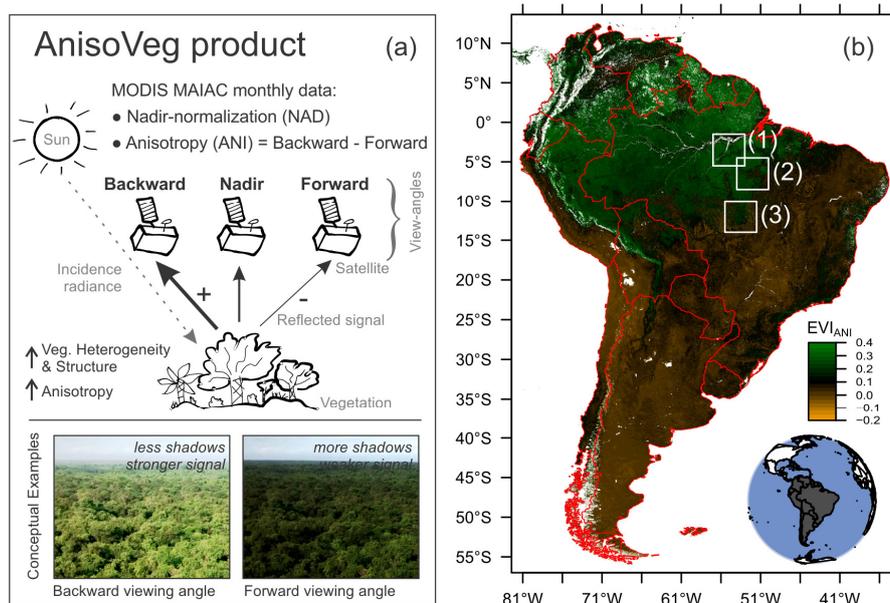
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143 **2. Methodology to compute the AnisoVeg product**

144 **2.1. Daily MODIS MAIAC surface reflectance data over South America**

145 Daily surface reflectance data were obtained from the MODIS product MCD19A1 v006
146 (collection 6) for the tiles covering South America (Figure 1). According to the MODIS traditional
147 tiling system, these tiles ranged from 9-14 (horizontal) and 7-14 (vertical). The input data
148 consisted in cross-calibrated surface reflectance from Terra and Aqua satellites on eight spectral
149 bands (Table 1) with 1-km spatial resolution from 2000 to 2021 (Lyapustin & Wang, 2018;
150 <http://dx.doi.org/10.5067/MODIS/MCD19A1.006>). This product provides surface reflectance
151 data corrected for atmospheric effects by the MAIAC algorithm, and controlled for cloud-free
152 and clear-to-moderately turbid conditions with Aerosol Optical Depth (AOD) at 0.47 μm below
153 1.5 (Lyapustin et al., 2018). The raw data were obtained from the NASA's Level-1 and
154 Atmosphere Archive and Distribution System (LAADS) Distributed Active Archive Center
155 (DAAC) available at <https://ladsweb.modaps.eosdis.nasa.gov/archive/allData/6/MCD19A1/>.

156



157

158 Figure 1 – AnisoVeg product concept and the area of coverage. (a) Schematic representation
 159 showing the observational geometry and the processing steps for producing NAD and ANI data
 160 from MODIS and to provide information on vegetation heterogeneity and structure, and (b) the
 161 visualization of the anisotropy EVI (EVI_{ANI}) for South America from August 2021 at 1-km spatial
 162 resolution, showing the coverage of the product in South America and the location of three sites
 163 used to demonstrate potential applications. The sites are: (1) Tapajós National Forest, (2) São
 164 Felix do Xingu, and (3) Xingu Park. Red lines indicate the countries boundaries.

165

166 Table 1 – MODIS spectral bands. NIR = near infrared; SWIR = shortwave infrared.

Band number	Band name	Wavelength (nm)
1	Red	620–670
2	NIR-1	841–876
3	Blue-1	459–479
4	Green	545–565
5	NIR-2	1230–1250
6	SWIR-1	1628–1652
7	SWIR-2	2105–2155
8	Blue-2	405–420

167

168 2.2. The AnisoVeg product

169 The AnisoVeg product consists of two types of data spanning from 2000 to 2021 in monthly
 170 composites at 1-km spatial resolution: (a) the nadir-normalized (NAD) data; and (b) the
 171 anisotropy (ANI) data. Each data type has 10 layers corresponding to the MODIS bands 1 to 8,
 172 and two VIs (NDVI and EVI).

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174



175 **2.2.1. The nadir-normalized (NAD) data**

176 In order to minimize the differences in sun-sensor geometry between the MODIS scenes and
 177 generate the NAD dataset, the daily surface reflectance data were normalized to a fixed 45° SZA
 178 and to nadir observation (VZA = 0°) using the BRDF and the Ross-Thick Li-Sparse (RTLS) model
 179 (Lucht and Lewis, 2000). Parameters of the RTLS BRDF model are part of the MAIAC product
 180 suite (MCD19A3 product) reported every 8 days. The closest RTLS parameters in time were used
 181 to normalize the daily data. The normalized Bidirectional Reflectance Factor (*BRFn*) for the NAD
 182 surface reflectance (SZA = 45°, VZA = 0°, RAA = 0°) was calculated using Eq. 1 (Lyapustin et
 183 al., 2018):

$$184 \quad BRFn = BRF \times \frac{k^L + F_{0V} \times k^V + F_{0G} \times k^G}{k^L + F_V \times k^V + F_G \times k^G} \quad (1)$$

185 where k^L , k^V , and k^G are the BRDF isotropic, volumetric, and geometric-optical kernel weights,
 186 respectively; F_{0V} and F_{0G} are the BRDF kernel values for the given geometry listed in Table 2;
 187 and F_V and F_G are the kernel values of the RTLS model for the specific MODIS observation,
 188 respectively (Lyapustin et al., 2018). F_V and F_G values are available at 5-km cells and were
 189 resampled to 1-km using the nearest neighbors' method to match the spatial resolution of the
 190 spectral bands. This resampling step does not create spatial artifacts in the data because the
 191 geometry changes slowly over time (Lyapustin et al., 2018).

192 Table 2 – View-angle normalizations and corresponding BRDF kernel values.

View-angle	F_{0V}	F_{0G}
Nadir	-0.04578	- 1.10003
Backward scattering	0.22930469	0.017440045
Forward scattering	-0.12029795	-1.6218740

193

194 We aggregated normalized daily data into monthly composites by keeping the median values for
 195 each pixel. During the temporal aggregation, we also calculated the per-pixel number of samples
 196 (or observations) for each monthly composite, which can be used as auxiliary data to filter pixels
 197 with low number of observations (less reliable estimates of surface reflectance). The tiles were
 198 mosaicked for the entire South America and then re-projected from the original sinusoidal
 199 projection to the geographic coordinates system (datum WGS-84, EPSG 4326). The output spatial
 200 resolution corresponded to 0.009107388 degrees, which is approximately equivalent to 1 km in
 201 projected coordinates.

202 We also calculated two traditional vegetation indices: NDVI (Rouse et al., 1973) (Eq. 2) and EVI
 203 (Huete et al., 2002) (Eq. 3).

$$204 \quad NDVI = \frac{\rho_{NIR} - \rho_{Red}}{\rho_{NIR} + \rho_{Red}} \quad (2)$$

$$205 \quad EVI = 2.5 * \frac{\rho_{NIR} - \rho_{Red}}{\rho_{NIR} + (6 * \rho_{Red} - 7.5 * \rho_{Blue}) + 1} \quad (3)$$

206 where ρ is the surface reflectance of a MODIS band, ρ_{NIR} is the NIR reflectance (band 2), ρ_{Red}
 207 is the red reflectance (band 1), and ρ_{Blue} is the blue reflectance (band 3). The constants in Eq. 3
 208 (6, 7.5, 1, and 2.5) represent: the aerosol coefficient adjustment of the atmosphere for the red and
 209 blue bands; the adjustment factor for the soil; and the gain factor, respectively (Huete et al., 2002).

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212



213 2.2.2. The anisotropy (ANI) data

214 For the ANI data, the daily surface reflectance data was first normalized to two viewing-angles at
215 the backward (SZA = 45°, VZA = 35°, RAA = 180°) and forward (SZA = 45°, VZA = 35°, RAA
216 = 0°) scattering using Eq. 1 and values from Table 2. To minimize potential errors of BRDF
217 extrapolation, the VZA was set to 35° instead of the hotspot (45°), because 35° is a very common
218 VZA in the empirical data distribution of the South America, and thus providing better estimates
219 of the anisotropy (Moura et al., 2015). The standard deviation for this modelling was thoroughly
220 investigated in a previous study and determined as 10% of the observed variation in anisotropy
221 (Moura et al., 2015). Further, we aggregated the backward and forward scattering data temporally
222 into monthly composites following the same procedures as before for the NAD data. We then
223 calculated the NDVI and EVI for each of the view-angle normalizations. Finally, we obtained the
224 difference between backward and forward scattering estimates for each of the eight MODIS
225 bands, as well as for the NDVI and EVI, effectively generating the ANI layers (Eq. 4; Moura et
226 al., 2015):

$$227 \text{ANI}_i = \text{Backward}_i - \text{Forward}_i \quad (4)$$

228 where i is the spectral band or VI selected in the calculation.

229

230 2.3. Algorithm and computation

231 All data processing was done in R v4.0.2 (R Core Team, 2016) and the code is available at GitHub
232 (https://github.com/ricds/maiac_processing) (Dalagnol & Wagner, 2022). Besides processing the
233 AnisoVeg product from the daily MAIAC MODIS data, the code can also generate 16-day or 8-
234 day temporal composites, mosaics, and VIs. Although we focused on South America when
235 developing AnisoVeg, the code can readily be adapted to process data for other parts of the world
236 and generate corresponding NAD and ANI layers. Below, we provide the computer specification
237 for anyone who wishes to process the data independently.

238 For the presented dataset, the computation was performed under a HP Z840 Workstation with
239 Intel Xeon CPU E5-2640 v3 (2.60Ghz, 32 cores), and 64 Gb RAM memory. The daily MODIS
240 data for the whole South America from 2000 to 2021 accounted for 6.69 Tb. Processing monthly
241 composites is computationally intensive due to loading all daily data for each month at once for
242 a given tile. Thus, the main bottlenecks are RAM memory and hard drive writing speed. For the
243 workstation with 64 Gb memory, the usage of 10 cores running in parallel processing was the
244 optimal choice. The average processing time of each monthly composite for one tile was 6
245 minutes. Therefore, it took 26.2 hours to process the 262 composites (March 2000 to December
246 2021) for each tile. Since we had 31 tiles covering the South America, the total amount of time to
247 process one view-normalization was approximately a month (33.8 days). Consequently, the total
248 time spent in computation was 101.5 days for processing the three view-normalizations (nadir,
249 backward, and forward scattering) and generating the NAD and ANI layers. Processing can also
250 be done with less potent computers with a minimum of 16 Gb RAM memory and 4 processing
251 cores.

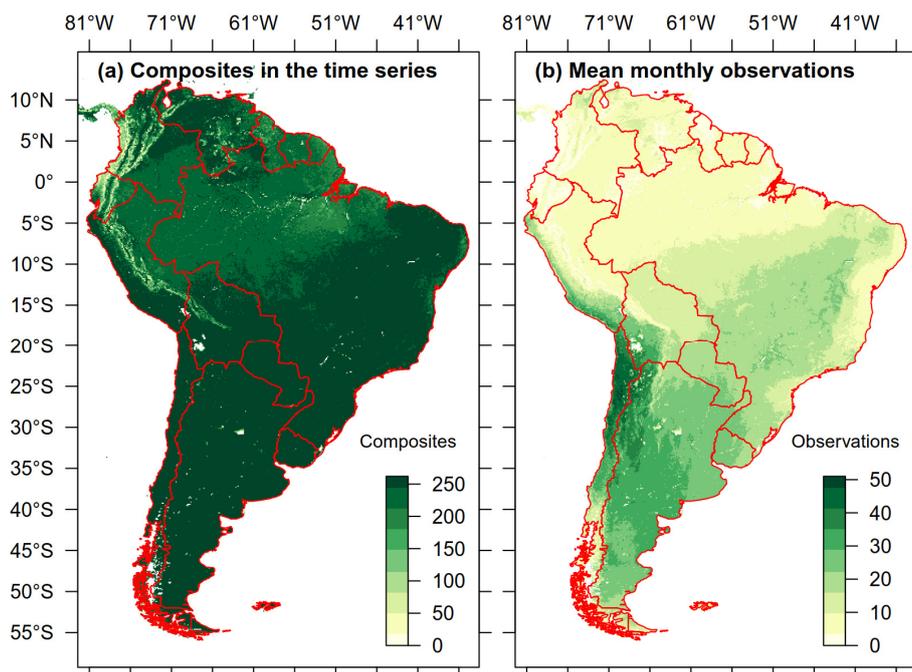
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253 2.4. Time series availability and uncertainty

254 The monthly compositing process returned a time series dataset over all of South America with
255 an average of 242 ± 35 out of a maximum of 262 composites (period between March 2000 and
256 December 2021) for each pixel with some data missing due to lack of high-quality observations
257 (Figure 2). Only 34.3% of the available pixels have the full time series (262 composites). The



258 Amazon region shows a lower mean number of samples in the time series with an average of 231
259 ± 29 composites, which can be seen in Figure 2. This lower number of samples is due to the innate
260 high cloud cover (Durieux et al., 2003). It is important to note that the AnisoVeg product was
261 strictly created to analyze land surface and does not cover water bodies. Moreover, the period
262 between March 2000 and June 2002 has higher amounts of missing data because it preceded the
263 launch of the Aqua satellite. When data from both satellites (Terra and Aqua) were combined to
264 create the product after 2002, we had a much better pixel level data availability to produce dense
265 time series. Although we have a dense time series across the Amazon rainforests (Figure 2a), the
266 mean number of daily observations within a month for this region is relatively lower than that
267 observed in more dry and seasonal regions of South America (Figure 2b). Thus, we suggest using
268 the number of samples layer as a proxy for uncertainty on the retrieval of monthly composites to
269 filter out pixels with low number of samples (e.g., less than three observations per composite).
270 The lesser number of samples one pixel has, the higher the uncertainty in the data analysis.



271

272 Figure 2 – AnisoVeg time series availability and uncertainty over South America. (a) The number
273 of composites in the time series representing pixel availability. The maximum number of
274 composites in the time series is 262 for the period between March 2000 and December 2021. (b)
275 Mean number of daily observations within a month used to create the monthly composites as a
276 proxy for uncertainty. The maximum daily observations in a composite are 60 (twice a day every
277 day for a month).

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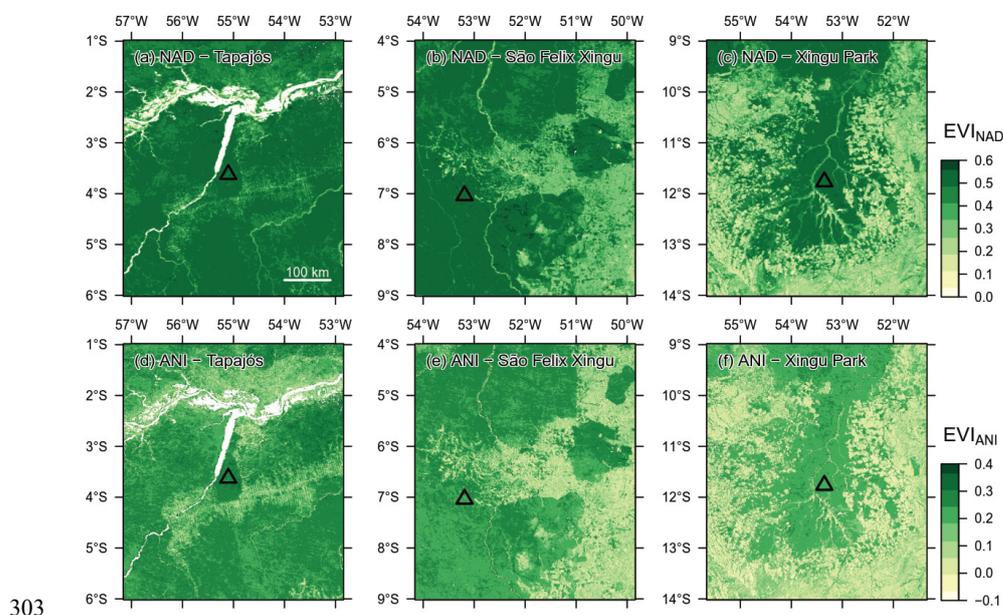
279 3. Spatial and temporal distribution of NAD and ANI data across the Amazon forests

280 To demonstrate the spatial and temporal distribution of NAD and ANI data over the Brazilian
281 Amazon rainforests, we selected three experimental areas (rectangles in Figure 1). These areas
282 show old-growth rainforests with distinct canopy structure and aboveground biomass (AGB)
283 stocks. The AGB increases from semideciduous forests at the Xingu Park ($190 \pm 19 \text{ Mg ha}^{-1}$) and



284 open ombrophilous forests with lianas at the São Felix do Xingu ($241 \pm 31 \text{ Mg ha}^{-1}$) to dense
285 ombrophilous forests at the Tapajós National Forest ($288 \pm 38 \text{ Mg ha}^{-1}$), as estimated by the
286 ESA/CCI AGB map from 2017 (Santoro & Cartus, 2021). These are large-scale AGB estimates
287 and may underestimate the true AGB at higher values such as in the Tapajós site. These three
288 sites are also expected to show different phenological dynamics because their selected pixels
289 cover distinct phenoregions in the study reported by Xu et al. (2015).

290 When compared to the nadir-normalized EVI (EVI_{NAD}) images (Figures 3a, b, c), the anisotropy
291 EVI (EVI_{ANI}) data showed different spatial patterns across sites (Figures 3d, e, f). While the
292 forests over the three sites showed approximately similar EVI_{NAD} values ($\text{EVI}_{\text{NAD}} \approx 0.50$) (Figures
293 3a,b,c), they showed more variability in EVI_{ANI} between the Xingu Park ($\text{EVI}_{\text{ANI}} > 0.20$), São
294 Felix do Xingu ($\text{EVI}_{\text{ANI}} > 0.24$), and Tapajós ($\text{EVI}_{\text{ANI}} > 0.27$) sites (Figures 3d,e,f). This increase
295 in EVI_{ANI} between sites goes into the same direction of the AGB gradient observed from the
296 Xingu Park to the Tapajós National Forest. This result may indicate different forest canopy
297 structures that were not captured in the EVI_{NAD} observations, but were captured by the EVI_{ANI} .
298 Overall, the EVI_{ANI} is high over forests (0.20 to 0.30) and low over pastures and crops (less than
299 0.10). This means large anisotropy between the reflected energy in backward and forward
300 scattering MODIS directions due to the structural complexity of forest canopies. The association
301 between anisotropy and forest canopy structure has been previously shown for the same region in
302 a previous work (Moura et al., 2016).

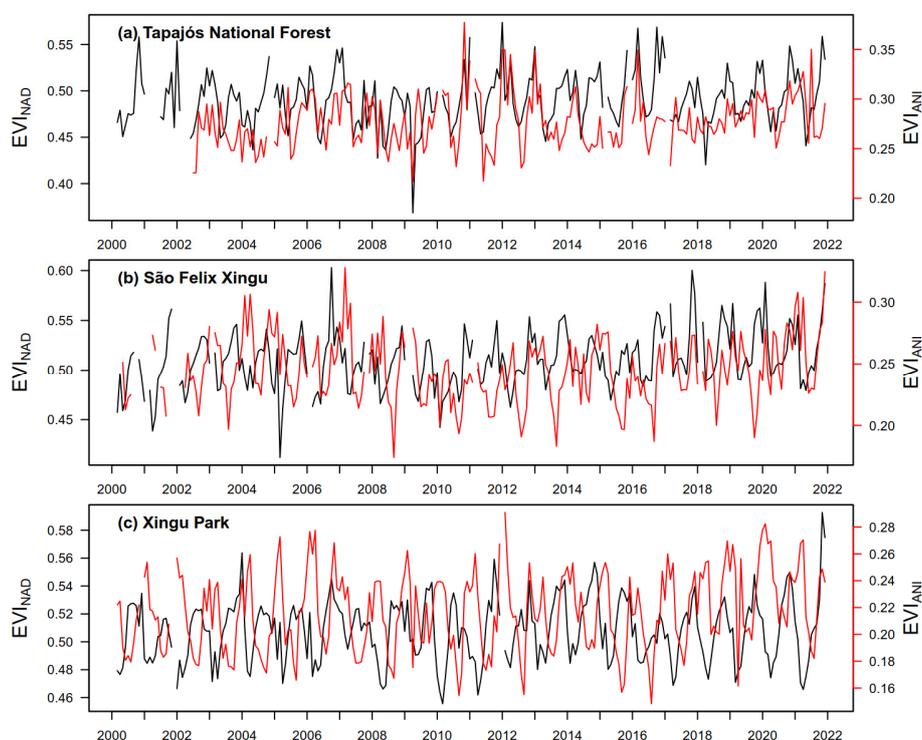


303
304 Figure 3 – The spatial distribution in August 2020 (dry season) of the nadir-normalized Enhanced
305 Vegetation Index (EVI_{NAD}) is shown in (a), (b), and (c) for the Tapajós National Forest, São Felix
306 do Xingu and Xingu Park, respectively. Corresponding results for the anisotropy EVI (EVI_{ANI})
307 are shown in (d), (e), and (f), respectively. The triangles plotted over (a, b, and c) indicate the
308 sites used to obtain the profiles of Figure 4.

309 From the comparison of different sites (triangles in Figure 3a), we observed that the mean EVI_{NAD}
310 signal over the time period did not vary much between the selected forests, while the EVI_{ANI}
311 varied greatly (Figure 4): Tapajós (mean $\text{EVI}_{\text{NAD}} = 0.49$, mean $\text{EVI}_{\text{ANI}} = 0.27$), São Felix do Xingu
312 (mean $\text{EVI}_{\text{NAD}} = 0.51$, mean $\text{EVI}_{\text{ANI}} = 0.24$), and Xingu Park (mean $\text{EVI}_{\text{NAD}} = 0.51$, mean EVI_{ANI}



313 = 0.22). Moreover, EVI_{NAD} and EVI_{ANI} values were moderately positively correlated at Tapajós
314 ($r = +0.37$), weakly correlated at São Felix do Xingu ($r = +0.06$), and moderately negatively
315 correlated at the Xingu Park ($r = -0.28$). The EVI_{NAD} and EVI_{ANI} are seasonal variability and phase
316 correlation changes from site to site, suggesting that different canopy dynamics processes are
317 likely being captured by the two metrics at the three sites. Understanding exactly what those
318 effects mean for these forests is beyond the scope of this paper. However, it indicates open venues
319 for studying forest functioning using these products. For example, previous studies have shown
320 that EVI_{NAD} metrics captured different compositions of leaf ages in the canopies of central
321 Amazon (Gonçalves et al., 2020).



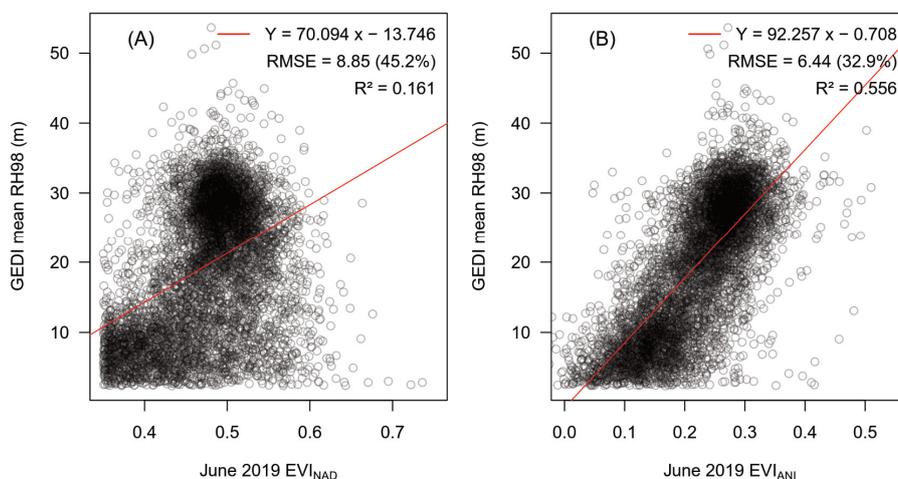
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323 Figure 4 – Time series of AnisoVeg’s MODIS Enhanced Vegetation Index (EVI) from 2000 to
324 2021 for old-growth forests of the (a) Tapajós National Forest; (b) São Felix do Xingu; (c) Xingu
325 Park. The black line indicates the nadir-normalized signal (NAD layer), while the red line
326 represents the EVI anisotropy (ANI layer). The profiles are the mean value of 3 x 3 pixels whose
327 locations are indicated by triangles in Figure 3.

328 To demonstrate the potential of AnisoVeg for large-scale forest structure inference, we compared
329 the NAD and ANI data against forest height measurements from the Global Ecosystem Dynamics
330 Investigation (GEDI) LiDAR sensor. We found that EVI_{ANI} was able to explain up to 55% of
331 height variability of Amazon forests according to a simple linear relationship ($R^2 = 0.55$, $p < 0.01$,
332 Figure 5). This is a very strong predicting power for a single variable, considering a simple linear
333 model, especially for satellite passive optical data which are often underrated for forest structure
334 estimates in comparison to Synthetic Aperture Radar (SAR) data. EVI_{NAD} was significantly but
335 weakly associated to height variability ($R^2 = 0.16$, $p < 0.01$), reinforcing the increase in
336 explanation power owed to the anisotropy metrics built from multi-angle observations. The height
337 data was derived from the GEDI LiDAR sensor aboard the International Space Station. They were



338 obtained more specifically from the product GEDI L2A elevation and height metrics data version
339 2 (footprint size 25 m), acquired from April 2019 to October 2020 (available dates at the time of
340 download). GEDI data were downloaded from Earth Data cloud service system
341 (<https://earthdata.nasa.gov>). We selected the Relative Height metric at 98th percentile (RH98),
342 which represents the top canopy height. The selected RH98 metric was averaged over each 1-km
343 grid cell, and filtered using a threshold of greater than or equal to 50 shots per km² to have a high
344 confidence of reliable height estimation representing the 1-km mean. The AnisoVeg data used for
345 this comparison were based on the same time period as GEDI, and filtered for EVI_{NAD} larger than
346 0.35 to exclude non-forested areas. While we only showed the plot for the strongest EVI_{ANI}:GEDI
347 relationship in June 2019 (Figure 5), the other months also showed significant ($p < 0.01$) and
348 strong relationships with R^2 ranging from 0.36 to 0.55 (mean $R^2 = 0.46$). Future studies should
349 explore relationships using ANI from different months and other indices, alone or in combination
350 with each other, to further understand their significance on explaining forest structure. This is
351 important to determine how the anisotropy data can contribute for aboveground biomass and
352 carbon estimates in conjunction with other sources of data such as those from SAR sensors.



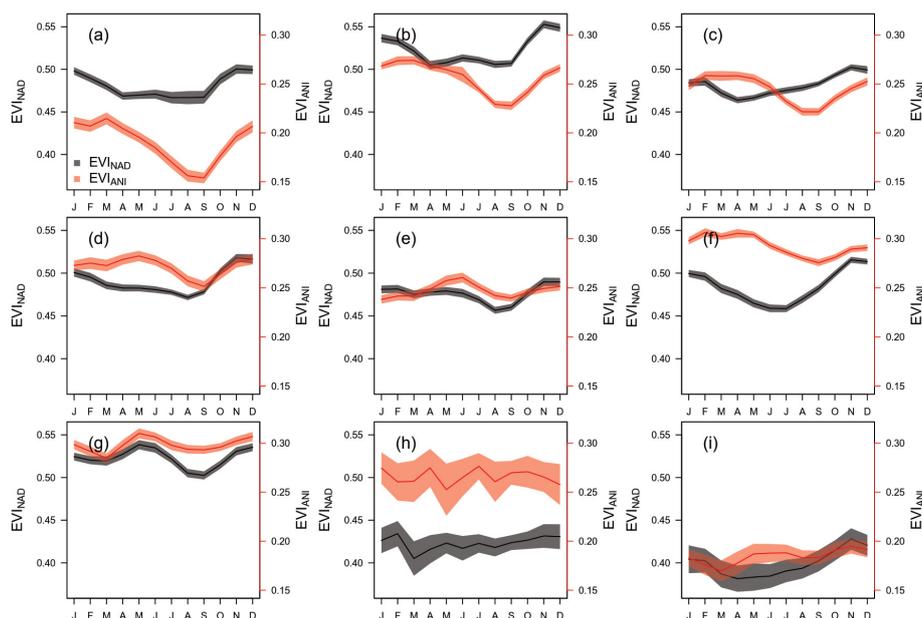
353

354 Figure 5 – Relationship between forest height (GEDi mean RH98) and two AnisoVeg layers
355 obtained in June 2019 over the Amazon: (a) EVI_{NAD} and (b) EVI_{ANI}. The RH98 metric consists in
356 the relative height at the 98th percentile, which represents the top of canopy height. 7,000 random
357 matching pixels were used in this analysis (1% of 700,000 total matching pixels available),
358 resulting from the filtering of both GEDI and AnisoVeg data. The red line indicates the fitted line
359 by a simple linear model.

360 In a prospective analysis, we also explored the behavior of the two EVI AnisoVeg metrics over
361 the Amazonian phenoregions mapped by Xu et al. (2015). The EVI_{NAD} and EVI_{ANI} monthly means
362 over different phenoregions highlighted the strong heterogeneity of the Amazonian forests
363 (Figure 6). For instance, the profiles showed strong differences between both metrics from
364 January to September in a phenoregion with well-defined dry and wet seasons (phenoregion one
365 in Figure 6a at the Xingu Park). Large differences between EVI_{NAD} and EVI_{ANI} were also observed
366 in some phenoregions without a very long dry season in northwest Amazon (phenoregion five in
367 Figure 6e). On the other hand, EVI_{NAD} and EVI_{ANI} showed temporal decoupling in phenoregion
368 three located at central-east Amazon (Figure 6c). Overall, while the seasonality of EVI_{NAD} has
369 been investigated by many studies in the past, the seasonality of EVI_{ANI} is something to be further
370 explored with the support of auxiliary data (e.g., airborne LiDAR and field campaigns). This is



371 important to better understand the differences in seasonal patterns between both AnisoVeg
372 metrics.



373
374 Figure 6 – Monthly means of EVI_{NAD} (black) and EVI_{ANI} (red) for nine phenoregions mapped by
375 Xu et al. (2015) in the Amazon. The phenoregions are shown in increasing order from 1 to 9 in
376 corresponding panels (a) to (i). They represent forests with similar seasonality and landscape
377 structure. Solid line and shaded area represent the mean and 95% confidence interval around the
378 mean. The values were extracted from 20 years of data (from 2001 to 2021) for 100 random
379 coordinates within each region, and extracted from 3 x 3 windows of pixels.

380

381 4. Prospective use of the dataset

382 The NAD layers from the AnisoVeg product have been used in previous studies to explore: the
383 climate drivers of the Amazon forest greening (Wagner et al., 2017); the large-scale Amazon
384 forest sensitivity to drought (Anderson et al., 2018); the structure and dominance of bamboo
385 species in southwest Amazon (Dalagnol et al., 2018); the productivity in a flooded forest in
386 eastern Amazon (Fonseca et al., 2019); the productivity and relationship with Sun-Induced
387 Fluorescence over the Brazilian Caatinga biome (Bontempo et al., 2020); the relationships with
388 leaf-age demography in central Amazon (Gonçalves et al., 2020); and the relationships with fire
389 disturbance and SAR-based Vegetation Optical Depth in southern Amazon (Zhang et al., 2021).

390 The ANI layers from the AnisoVeg product have been mainly used to characterize Amazon forest
391 structure properties (Moura et al., 2015; 2016). These layers now open new venues of
392 investigation on vegetation, including (but not limited to): the characterization of biophysical
393 attributes of forests, including their seasonality and trends; the assessment of changes in
394 vegetation structure due to natural disturbances or degradation (logging, fire, edge effects); and
395 the evaluation of forest health and productivity (greenness and browning). We expect that this
396 dataset contributes to upscaling studies over large areas of key forest properties such as the AGB
397 and canopy roughness (Foody & Curran, 1994; Saatchi et al., 2008). This information is required
398 for dynamic vegetation models to accurately represent the carbon cycle. This dataset is not limited



399 to study Amazonian forests and can be used to explore other biomes of South America such as
 400 the Atlantic Forest, savannas (Cerrado), Caatinga, Chaco, Pantanal, and Pampas. Such studies
 401 could improve our understanding of large-scale vegetation functioning, carbon storage, and
 402 cycling. Ultimately, they can contribute to refine global ecosystem models, and to obtain accurate
 403 estimates of carbon cycle in response to climate and environmental change. Furthermore,
 404 auxiliary backward and forward scattering data are also available with the dataset. Beyond the
 405 use of the provided ANI layers, this effectively allows the computation of several other multi-
 406 angular anisotropy indices from the literature (Table 3), offering the possibility to investigate their
 407 use for tropical vegetation studies.

408

409 Table 3 – Examples of other multi-angular anisotropy indices that can be further calculated using
 410 layers of the AnisoVeg product. Lambda represents the selected spectral band or vegetation index.
 411 N, H, and D represent nadir-view normalization, hot-spot (backward scattering), and dark-spot
 412 (forward scattering) estimates, respectively.

Anisotropy Indices	Formula	Reference
Anisotropy index (ANIX)	$\frac{\lambda_H}{\lambda_D}$	Sandmeier et al. (1998)
Nadir BRDF-adjusted NDVI (NDVI_{ISO})	$\frac{NIR_N - RED_N}{NIR_N + RED_N}$	Schaaf et al. (2002)
Hot-spot dark-spot index (HDS_{RED})	$\frac{RED_H - RED_D}{RED_D}$	Lacaze et al. (2002)
Normalized difference between hot-spot and dark-spot index (NDHD_{NIR})	$\frac{NIR_H - NIR_D}{NIR_H + NIR_D}$	Chen et al. (2005)
Hot-spot dark-spot NDVI (NDVI_{HD})	$\frac{NIR_H - RED_D}{NIR_H + RED_D}$	Pocewicz et al. (2007)
Hot-spot-incorporated NDVI (NDVI_{HS})	$NDVI_N \times (1 - RED_H)$	Pocewicz et al. (2007)
Anisotropy difference (ANI)*	$\lambda_H - \lambda_D$	Moura et al. (2015)
Vegetation Structure Index (VSI)	$\frac{NDVI_D - NDVI_H}{1 - NIR_D}$	Sharma et al. (2021)

413 *ANI is included in the AnisoVeg product. Source: Adapted from Sharma et al. (2021).

414

415 5. Code and data availability

416 All code is available at GitHub (https://github.com/ricds/maiac_processing) (Dalagnol &
 417 Wagner, 2022). The full dataset can be found at the official AnisoVeg repository at Zenodo
 418 (<https://doi.org/10.5281/zenodo.3878879>) (Dalagnol et al., 2022). The dataset was organized in
 419 compressed files (“.zip” format) sub-divided by years (currently 2000-2021) and layers (bands 1-
 420 8, NDVI, and EVI) for both nadir-normalization (code = NAD) and anisotropy (code = ANI). The
 421 number of samples layers (code = NO_SAMPLES) are also provided. Inside each compressed
 422 file there will be 12 image files (“.tif” format), one per month, except for the year 2000 which
 423 starts in March. The storage size for the whole dataset is 162.6 Gb. The data have a scale factor
 424 of 10,000 to reduce file storage size. Thus, to obtain surface reflectance values of bands or correct
 425 range of values for indices, you should divide the layers by 10,000. The exception is the number
 426 of samples, which already shows the correct range of values from 0 to 60 observations. The dataset
 427 is planned to be updated on a yearly-basis. Auxiliary data that allow the calculation of other
 428 anisotropy metrics (listed in Table 3) are included in two separate Zenodo repositories for
 429 backward (<https://doi.org/10.5281/zenodo.6040300>) (Dalagnol, 2022a) and forward scattering



430 (<https://doi.org/10.5281/10.5281/zenodo.6048785>) (Dalagnol, 2022b), including the selected
431 layers Red, NIR, NDVI and EVI. The EVI_{ANI} and EVI_{NAD} layers were also uploaded to the GEE
432 platform using the *geeup* tool v0.5.3 (Roy, 2022). They can be accessed through the GEE
433 ImageCollection assets “projects/anisoveg/assets/evi_anisotropy” and
434 “projects/anisoveg/assets/evi_nadir”, found at
435 https://code.earthengine.google.com/?asset=projects/anisoveg/assets/evi_anisotropy and
436 https://code.earthengine.google.com/?asset=projects/anisoveg/assets/evi_nadir.

437

438 **Author contribution**

439 R.D. and Y.M. conceived the presented idea. R.D. designed the methodology with contributions
440 from Y.M. on the anisotropy method. R.D. conducted formal analysis and investigation with
441 contributions from L.G., F.W., N.G., and S.S. Y.W. and A.L. provided the original MODIS
442 (MAIAC) data and support for processing it. Y.Y. and S.S. provided the processed GEDI height
443 data and support to analyze it. R.D. and F.W. developed the code to process the MODIS (MAIAC)
444 data into the products. R.D. conducted data curation of the products. L.A. supervised the project.
445 R.D. wrote the original draft with support from L.G., F.W. and Y.M. All authors read, reviewed
446 and approved the final version of the manuscript.

447

448 **Acknowledgements**

449 R.D. was supported by Sao Paulo Research Foundation (FAPESP) grants 2015/22987-7 and
450 2019/21662-8. F.W. was supported by FAPESP grant 2015/50484-0. Part of this work was carried
451 out at the Jet Propulsion Laboratory, California Institute of Technology, under a contract with the
452 National Aeronautics and Space Administration (NASA). The funders had no role in the study
453 design, data collection and analysis, including the decision to publish or prepare the manuscript.
454 We thank the MODIS MAIAC team from NASA for providing the freely available MODIS
455 (MAIAC) daily dataset.

456

457 **Conflict of Interest**

458 The authors have declared no conflict of interest.

459

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