## 1 AnisoVeg: Anisotropy and Nadir-normalized MODIS MAIAC datasets for satellite 2 vegetation studies in South America

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- 26 This document includes: main manuscript, figures and tables.

#### 27 Abstract

28 The AnisoVeg product consists of monthly 1-km composites of anisotropy (ANI) and nadir-29 normalized (NAD) surface reflectance layers obtained from the Moderate Resolution Imaging 30 Spectroradiometer (MODIS) sensor over the entire South America. The satellite data were pre-31 processed using the Multi-Angle Implementation Atmospheric Correction (MAIAC). The 32 AnisoVeg product spans 22 years of observations (2000 to 2021) and includes the reflectance of 33 MODIS bands 1 to 8 and two vegetation indices (VIs): Normalized Difference Vegetation Index 34 (NDVI) and Enhanced Vegetation Index (EVI). While the NAD layers reduce the data variability 35 added by bidirectional effects on the reflectance and VI time series, the unique ANI layers allow 36 the use of this multi-angular data variability as a source of information for vegetation studies. The 37 AnisoVeg product has been generated using daily MODIS MAIAC data from both Terra and 38 Aqua satellites, normalized for a fixed solar zenith angle (SZA =  $45^{\circ}$ ), modelled for three sensor 39 view directions (nadir, forward, and backward scattering), and aggregated to monthly composites. 40 The anisotropy was calculated by the subtraction of modelled backward and forward scattering 41 surface reflectance. The release of the ANI data for open usage is novel, as well as the NAD data 42 at an advance processing level. We demonstrate the use of such data for vegetation studies using 43 three types of forests in eastern Amazon with distinct gradients of vegetation structure and 44 aboveground biomass (AGB). The gradient of AGB was positively associated with ANI, while 45 NAD values were related to different canopy structural characteristics. This was further illustrated by the strong and significant relationship between EVI<sub>ANI</sub> and forest height observations from the 46 Global Ecosystem Dynamics Investigation (GEDI) LiDAR sensor considering a simple linear 47 48 model ( $R^2 = 0.55$ ). Overall, the time series of the AnisoVeg product (NAD and ANI) provide 49 distinct information for various applications aiming at understanding vegetation structure, 50 dynamics, and disturbance patterns. All data, processing codes and results are made publicly 51 available to enable research and the extension of AnisoVeg products for other regions outside the 52 South America. The code can be found at https://doi.org/10.5281/zenodo.6561351 (Dalagnol and 53 Wagner, 2022),  $EVI_{ANI}$  and  $EVI_{NAD}$  can be found as assets in the Google Earth Engine (GEE) 54 (described in the data availability section), and the full dataset is available at the open repository 55 <a href="https://doi.org/10.5281/zenodo.3878879">https://doi.org/10.5281/zenodo.3878879</a>> (Dalagnol et al., 2022).

56 Key-words: AnisoVeg, South America, vegetation structure, forest monitoring, MODIS.

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### 58 1. Introduction

59 The anisotropy is defined as the departure from Lambertian scattering (isotropic), caused by the 60 physical structure of media through which photons pass. The anisotropy is defined by the 61 directional dependence of observations on mechanical or physical properties of surfaces. Because 62 most land covers are not Lambertian (isotropic), the surface reflectance measured by satellite sensors varies with the view zenith angle (VZA), view direction (backward or forward scattering), 63 64 and solar zenith angle (SZA) (Galvão et al., 2011). This is especially valid for images acquired 65 over vegetated surfaces by large field-of-view (FOV) instruments such as the Moderate 66 Resolution Imaging Spectroradiometer (MODIS) (Bhandari et al., 2011). MODIS has a wide swath scanning  $\pm 55^{\circ}$  from nadir on board the Terra and Aqua satellites. For example, a reflected 67 68 signal coming from the backward scattering direction of MODIS under a large VZA and close-69 to-zero relative azimuth angle (RAA) between the satellite and sun (sun behind the platform) is 70 generally higher than that coming from the nadir (VZA =  $0^{\circ}$ ) or forward scattering direction 71 (platform facing the sun at  $RAA = 180^{\circ}$ ). Moreover, the SZA also varies seasonally and across 72 geographical locations, affecting the amounts of shadows in the surfaces observed by satellites 73 (Galvão et al., 2013). Such view-illumination effects are dependent on the land cover types and 74 their magnitude relates to differences in biophysical properties of the vegetation (Galvão et al., 75 2004; Sims et al., 2011Foody & Curran, 1994). Therefore, the vegetation anisotropy can be seen 76 antagonistically as sources of noise and biophysical information in the time-series analysis of 77 vegetation indices (VIs) calculated from MODIS. As a source of noise, one may consider that the 78 reflected signal toward the large FOV satellite sensors varies with distinct view-illumination 79 geometries of data acquisition over the same surface. As a source of information, one may 78 highlight that the anisotropy is land-cover type dependent, showing spectral variations that may 80 be associated, for instance, with changes in vegetation structure across different forests.

82 To reduce the bidirectional effects as a source of noise, a nadir-normalized dataset can be created. 83 We can normalize the surface reflectance of the MODIS bands to a specific set of VZA and SZA 84 using the bidirectional reflectance distribution function (BRDF), represented by a model such as 85 the Ross-Thick Li-Sparse (RTLS) (Wanner et al., 1995). To ensure confidence in the data 86 analysis, we can also use the Multi-Angle Implementation Atmospheric Correction (MAIAC) for 87 atmospheric correction. MAIAC is a new generation of cloud screening and atmospheric 88 correction algorithm that uses an adaptive time series analysis and processing of groups of pixels 89 to derive atmospheric aerosol concentration, cloud mask and surface reflectance without typical 90 empirical assumptions (Lyapustin et al., 2011, 2012). By mitigating atmospheric interference and 91 advancing the accuracy of surface reflectance over tropical vegetation by a factor of 3 to 10, 92 MAIAC offers substantial improvement over conventional products such as the MOD09 (Hilker 93 et al., 2012). Because of the better data quality retrieval, MAIAC is also an alternative to the 94 MCD43A4 16-day Nadir Bidirectional Reflectance Distribution Function (BRDF)-Adjusted 95 Reflectance (NBAR) product due to the less variable seasonal signal (3 to 10 times) over 96 evergreen forests resultant from reduced effects of sun-view geometry. While the MCD43A4 97 NBAR product offers view-illumination correction, using the MAIAC products one can also 98 correct for solar illumination effects at the same time. It offers substantial improvement over 99 conventional algorithms by mitigating atmospheric interference and advancing the accuracy of 100 surface reflectance over tropical vegetation by a factor of 3 to 10 (Hilker et al., 2012). Due to the 101 improvements in cloud detection, aerosol retrieval and atmospheric correction, the MAIAC 102 algorithm provides from 4 to 25% more high-quality retrievals than the traditional MOD09 103 product, with the largest estimate being observed for tropical regions (Lyapustin et al., 2021). 104 Studies have used MODIS MAIAC observations with nadir-normalized geometry to assess 105 Amazonian forests' structure, functioning, and impacts of environmental and climate change 106 (Hilker et al., 2014; Wagner et al., 2017; Anderson et al., 2018; Dalagnol et al., 2018; Fonseca et 107 al., 2019; Bontempo et al., 2020; Goncalves et al., 2020; Zhang et al., 2021). For instance, such 108 products provided reliable time series of surface reflectance data that allowed to identify large-109 scale communities of bamboo species and their dynamics in the southwest Amazon (Dalagnol et 110 al., 2018). Lastly, by improving the cloud screening and minimizing BRDF artifacts in 111 comparison to uncorrected data, the MAIAC greatly contributed to the understanding of the long-112 standing debate in the Amazon over the possible existence of the green-up phenomenon observed 113 during the dry season of each year or with severe droughts (Morton et al., 2014; Bi et al., 2015; 114 Saleska et al., 2016; Wu et al., 2017). The existence of this phenomenon has implications on the 115 comprehension of the resilience of tropical forests to climate change.

116 To use the bidirectional effects as a source of information, we generate an anisotropy dataset that 117 is dependent on land-cover types and captures the variations of sunlit and shaded canopy 118 components viewed by the sensors (Chen et al., 2003; Gao, 2003). The use of multi-angular 119 information to obtain metrics of anisotropy and extract information on forest structure was 120 suggested two decades ago (Gobron et al., 2002; Diner et al., 2005Foody & Curran, 1994). One 121 of the early experiments exploring the use of anisotropy to extract information about vegetation structure were conducted by calculating the ratio between backward and forward scattering data 122 and generating the anisotropy index (ANIX) on studying short-stature grass-type vegetation 123 124 (Sandmeier et al., 1998). The first experiments with such concept were conducted by calculating 125 the ratio between backward and forward scattering data and generating the anisotropy index 126 (ANIX) on studying short stature grass type vegetation (Sandmeier et al., 1998). Other indices 127 have been developed and validated afterwards (Schaaf et al., 2002; Lacaze et al., 2002; Chen et 128 al., 2005; Pocewicz et al., 2007; Moura et al., 2015; Sharma et al., 2021). However, this remains 129 an understudied topic with limited results reported in the literature, especially in tropical regions. 130 For instance, observations from the Multi-angle Imaging Spectroradiometer (MISR)/Terra in the 131 backward and forward scattering directions facilitated the discrimination of savanna physiognomies in Brazil (Liesenberg et al., 2007). MODIS MAIAC data from both directions 132 133 were also used to calculate an anisotropic VI that explained part of the large-scale photosynthetic activity in the Amazon, where higher photosynthetic activity was associated to higher anisotropy 134 135 values (Sousa et al., 2017). Moura et al. (2015) employed a more sophisticated approach based 136 on scattering at backward and forward view directions using multi-temporal and multi-angular 137 observations of MAIAC MODIS and BRDF modelling. The resultant metrics of anisotropy were 138 further validated against field and airborne Light Detection And Ranging (LiDAR) observations, showing strong linear relationship with leaf area index (LAI) ( $R^2 = 0.70-0.88$ ), canopy 139 140 heterogeneity ( $R^2 = 0.54$ ), and photosynthetic activity ( $R^2 = 0.73 - 0.98$ ) (Moura et al., 2015; Moura 141 et al., 2016; Hilker et al., 2017). Although showing great potential in vegetation studies, the 142 aforementioned anisotropy metrics were never computed over larger areas of the world such as 143 proposed in this study for South America.

144 The objective of this work is to present the AnisoVeg product, and how it can be used for 145 vegetation studies. We use MODIS Collection 6 (C6) MAIAC (Lyapustin et al., 2018) monthly 146 data (2000-2021) generated at 1-km spatial resolution for the entire South America with two 147 different types of layers: (1) nadir-normalized (NAD) data for the surface reflectance of MODIS 148 bands 1 to 8 and two VIs (NDVI and EVI); and (2) anisotropy data (ANI) calculated from the 149 difference between backward and forwarding scattering estimates of bands 1 to 8 and VIs (Moura 150 et al., 2015). The motivations for generating this product extend from developing applications of 151 multi-angle observations for vegetation studies to producing analysis-ready and openly available 152 datasets of anisotropy and nadir metrics for a larger community of users. The paper is organized 153 in several sections to present the processing steps for generating the AnisoVeg products, a brief 154 evaluation of data products over experimental areas, and finally an example of its potential 155 application in vegetation studies.

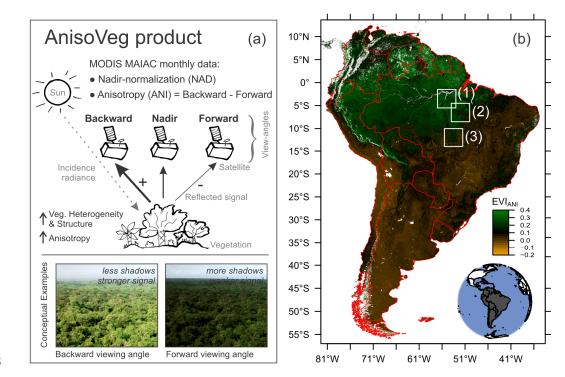
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# 157 **2. Methodology to compute the AnisoVeg product**

# 158 **2.1. Daily MODIS MAIAC surface reflectance data over South America**

159 Daily surface reflectance data were obtained from the MODIS product MCD19A1 v006 160 (collection 6) for the tiles covering South America (Figure 1). According to the MODIS traditional tiling system, these tiles ranged from 9-14 (horizontal) and 7-14 (vertical). The input data 161 162 consisted in cross-calibrated surface reflectance from Terra and Aqua satellites on eight spectral bands (Table 1) with 1-km spatial resolution from 2000 to 2021 (Lyapustin & Wang, 2018; 163 164 http://dx.doi.org/10.5067/MODIS/MCD19A1.006). This product provides surface reflectance 165 data corrected for atmospheric effects by the MAIAC algorithm, and controlled for cloud-free 166 and clear-to-moderately turbid conditions with Aerosol Optical Depth (AOD) at 0.47 µm below 167 1.5 (Lyapustin et al., 2018). The MAIAC algorithm uses a time series approach for improved 168 cloud filtering amongst other filters such as surface reflectance change in order to provide the 169 most accurate surface reflectance estimates. The raw data were obtained from the NASA's Level-170 1 and Atmosphere Archive and Distribution System (LAADS) Distributed Active Archive Center

171 (DAAC) available at https://ladsweb.modaps.eosdis.nasa.gov/archive/allData/6/MCD19A1/.



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Figure 1 – AnisoVeg product concept and the area of coverage. (a) Schematic representation showing the observational geometry and the processing steps for producing NAD and ANI data from MODIS and to provide information on vegetation heterogeneity and structure, and (b) the visualization of the anisotropy EVI (EVI<sub>ANI</sub>) for South America from August 2021 at 1-km spatial resolution, showing the coverage of the product in South America and the location of three sites used to demonstrate potential applications. The sites are: (1) Tapajós National Forest, (2) São Felix do Xingu, and (3) Xingu Park. Red lines indicate the countries boundaries.

181

182 Table 1 – MODIS spectral bands. NIR = near infrared; SWIR = shortwave infrared.

Band number	Band name	Wavelength (nm)
1	Red	620–670
2	NIR-1	841-876
3	Blue-1	459–479
4	Green	545-565
5	NIR-2	1230-1250
6	SWIR-1	1628–1652
7	SWIR-2	2105-2155
8	Blue-2	405–420

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## 184 2.2. The AnisoVeg product

185 The AnisoVeg product consists of two main types of data spanning from 2000 to 2021 in monthly

186 composites at 1-km spatial resolution: (a) the nadir-normalized (NAD) data; and (b) the

187 anisotropy (ANI) data. Each data type has 10 layers corresponding to the MODIS bands 1 to 8,

188 and two VIs (NDVI and EVI). Additionally, the product provides auxiliary layers of backward

189 scattering and forward scattering, including part of the bands (description on section 5). The

190 AnisoVeg product consists of two types of data spanning from 2000 to 2021 in monthly 191 composites at 1 km spatial resolution: (a) the nadir normalized (NAD) data; and (b) the

anisotropy (ANI) data. Each data type has 10 layers corresponding to the MODIS bands 1 to 8,

193 and two VIs (NDVI and EVI).

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### 196 2.2.1. The nadir-normalized (NAD) data

197 In order to minimize the differences in sun-sensor geometry between the MODIS scenes and 198 generate the NAD dataset, the daily surface reflectance data were normalized to a fixed 45° SZA 199 and to nadir observation (VZA =  $0^{\circ}$ ) using the BRDF and the Ross-Thick Li-Sparse (RTLS) model 200 (Wanner et al., 1995Lucht and Lewis, 2000). Parameters of the RTLS BRDF model are part of 201 the MAIAC product suite (MCD19A3 product) reported every 8 days. The MAIAC algorithm 202 detects significant land cover changes (e.g. fire, deforestation) within the 8-day period and does 203 not use those observations for the BRDF inversion (Lyapustin et al., 2018). A minimum of three 204 observations in the eight-day window was required to accurately model the signal. The closest RTLS parameters in time were used to normalize the daily data. The normalized Bidirectional 205 206 Reflectance Factor (*BRFn*) for the NAD surface reflectance (SZA =  $45^\circ$ , VZA =  $0^\circ$ , RAA =  $0^\circ$ ) 207 was calculated using Eq. 1 (Lyapustin et al., 2018):

$$208 \qquad BRFn = BRF \times \frac{k^L + F_{0V} \times k^V + F_{0G} \times k^G}{k^L + F_V \times k^V + F_G \times k^G}$$
(1)

where  $k^L$ ,  $k^V$ , and  $k^G$  are the BRDF isotropic, volumetric, and geometric-optical kernel weights, respectively;  $F_{0V}$  and  $F_{0G}$  are the BRDF kernel values for the given geometry listed in Table 2; and  $F_V$  and  $F_G$  are the kernel values of the RTLS model for the specific MODIS observation, respectively (Lyapustin et al., 2018).  $F_V$  and  $F_G$  values are available at 5-km cells and were resampled to 1-km using the nearest neighbors' method to match the spatial resolution of the spectral bands. This resampling step does not create spatial artifacts in the data because the geometry changes slowly over time (Lyapustin et al., 2018).

View-angle	<u>Solar</u> Zenith	<u>View</u> Zenith	<u>Relative</u> Azimuth	F <sub>0V</sub>	$F_{\theta G}$
	Angle	Angle	Angle		
	$(\overline{SZA, o})$	$(VZA, ^{o})$	$(RAA, ^{o})$		
Nadir	45	0	0	-0.04578	- 1.10003
Backward scattering	<u>45</u>	<u>35</u>	180	0.22930469	0.01744004
Forward scattering	45	35	0	-0.12029795	-1.6218740

216 Table 2 – View-angle normalizations and corresponding BRDF kernel values.

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218 We aggregated normalized daily data into monthly composites by keeping the median values for 219 each pixel. During the temporal aggregation, we also calculated the per-pixel number of samples 220 (or observations) for each monthly composite, which can be used as auxiliary data to filter pixels 221 with low number of observations (less reliable estimates of surface reflectance). The tiles were 222 mosaicked for the entire South America and then re-projected from the original sinusoidal 223 projection to the geographic coordinates system (datum WGS-84, EPSG 4326). The output spatial 224 resolution corresponded to 0.009107388 degrees, which is approximately equivalent to 1 km in 225 projected coordinates.

We also calculated two traditional vegetation indices: NDVI (Rouse et al., 1973) (Eq. 2) and EVI
(Huete et al., 2002) (Eq. 3).

228 
$$NDVI = \frac{\rho NIR - \rho Red}{\rho NIR + \rho Red}$$
 (2)

$$229 \quad EVI = 2.5 \times * \frac{\rho NIR - \rho Red}{\rho NIR + (6 \times * \rho Red - 7.5 \times * \rho Blue) + 1}$$
(3)

where  $\rho$  is the surface reflectance of a MODIS band,  $\rho NIR$  is the NIR reflectance (band 2),  $\rho Red$ is the red reflectance (band 1), and  $\rho Blue$  is the blue reflectance (band 3). The constants in Eq. 3 (6, 7.5, 1, and 2.5) represent: the aerosol coefficient adjustment of the atmosphere for the red and blue bands; the adjustment factor for the soil; and the gain factor, respectively (Huete et al., 2002).

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### 237 2.2.2. The anisotropy (ANI) data

238 For the ANI data, the daily surface reflectance data was first normalized to two viewing-angles at 239 the backward (SZA =  $45^\circ$ , VZA =  $35^\circ$ , RAA =  $180^\circ$ ) and forward (SZA =  $45^\circ$ , VZA =  $35^\circ$ , RAA 240  $= 0^{\circ}$ ) scattering using Eq. 1 and values from Table 2. The VZA was set to near hotspot (VZA = 241  $35^{\circ}$ ) instead of the actual hotspot (VZA =  $45^{\circ}$ ) to keep VZA closer to the actual range of MODIS 242 observations across the South America and minimize errors coming from extrapolation of the 243 BRDF (Moura et al., 2015). To minimize potential errors of BRDF extrapolation, the VZA was 244 set to 35° instead of the hotspot (45°), because 35° is a very common VZA in the empirical data 245 distribution of the South America, and thus providing better estimates of the anisotropy (Moura et al., 2015). The standard deviation for this modelling was thoroughly investigated in a previous 246 247 study and determined as 10% of the observed variation in anisotropy (Moura et al., 2015). Further, 248 we aggregated the backward and forward scattering data temporally into monthly composites 249 following the same procedures as before for the NAD data. We then calculated the NDVI and 250 EVI for each of the view-angle normalizations. Finally, we obtained the difference between 251 backward and forward scattering estimates for each of the eight MODIS bands, as well as for the 252 NDVI and EVI, effectively generating the ANI layers (Eq. 4; Moura et al., 2015):

$$253 \quad ANI_i = Backward_i - Forward_i$$

(4)

- where i is the spectral band or VI selected in the calculation.
- 255

#### 256 **2.3. Algorithm and computation**

All data processing was done in R v4.0.2 (R Core Team, 2016) and the code is available at GitHub (https://github.com/ricds/maiac\_processing) (Dalagnol & Wagner, 2022). Besides processing the AnisoVeg product from the daily MAIAC MODIS data, the code can also generate 16-day or 8day temporal composites, mosaics, and VIs. Although we focused on South America when developing AnisoVeg, the code can readily be adapted to process data for other parts of the world and generate corresponding NAD and ANI layers. Below, we provide the computer specification for anyone who wishes to process the data independently.

For the presented dataset, the computation was performed under a HP Z840 Workstation with Intel Xeon CPU E5-2640 v3 (2.60Ghz, 32 cores), and 64 Gb-GB (gigabytes) RAM memory. The daily MODIS data for the whole South America from 2000 to 2021 accounted for 6.69 TbTB 267 (terabytes). Processing monthly composites is computationally intensive due to loading all daily 268 data for each month at once for a given tile. Thus, the main bottlenecks are RAM memory and 269 hard drive writing speed. For the workstation with 64 Gb-GB memory, the usage of 10 cores 270 running in parallel processing was the optimal choice. The average processing time of each 271 monthly composite for one tile was 6 minutes. Therefore, it took 26.2 hours to process the 262 272 composites (March 2000 to December 2021) for each tile. Since we had 31 tiles covering the 273 South America, the total amount of time to process one view-normalization was approximately a 274 month (33.8 days). Consequently, the total time spent in computation was 101.5 days for 275 processing the three view-normalizations (nadir, backward, and forward scattering) and 276 generating the NAD and ANI layers. Processing can also be done with less potent computers with 277 a minimum of 16 Gb-GB RAM memory and 4 processing cores.

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### 279 2.4. Time series availability and uncertainty

280 The monthly compositing process returned a time series dataset over all of South America with 281 an average of  $242 \pm 35$  out of a maximum of 262 composites (period between March 2000 and 282 December 2021) for each pixel with some data missing due to lack of high-quality observations 283 (Figure 2). Only 34.3% of the available pixels have the full time series (262 composites). The 284 Amazon region shows a lower mean number of samples in the time series with an average of 231 285  $\pm$  29 composites, which can be seen in Figure 2. This lower number of samples is due to the innate high cloud cover (Durieux et al., 2003). It is important to note that the AnisoVeg product was 286 287 strictly created to analyze land surface and does not cover water bodies. Moreover, the period 288 between March 2000 and June 2002 has higher amounts of missing data because it preceded the 289 launch of the Aqua satellite. When data from both satellites (Terra and Aqua) were combined to 290 create the product after 2002, we had a much better pixel level data availability to produce dense 291 time series. Although we have a dense time series across the Amazon rainforests (Figure 2a), the 292 mean number of daily observations within a month for this region is relatively lower than that 293 observed in more dry and seasonal regions of South America (Figure 2b). Thus, we suggest using 294 the number of samples layer as a proxy for uncertainty on the retrieval of monthly composites to 295 filter out pixels with low number of samples (e.g., less than three observations per composite). 296 The lesser number of samples one pixel has, the higher the uncertainty in the data analysis. 297 Although we use the median values to aggregate observations within months and mitigate 298 potential land cover changes, stand-replacing changes may cause inaccurate anisotropy estimates 299 for the given monthly estimates. Hence, we advise filtering data for land use and land cover 300 changes before using them to obtain the most accurate anisotropy estimates.

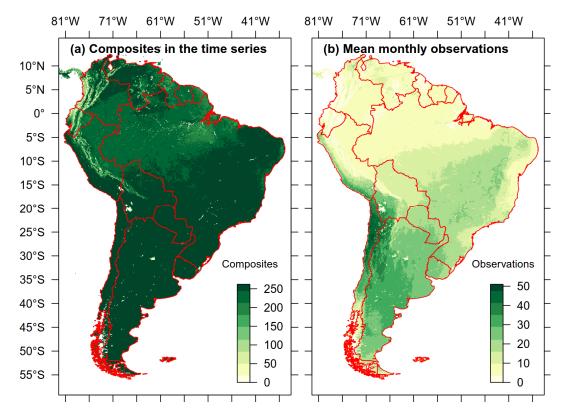




Figure 2 – AnisoVeg time series availability and uncertainty over South America. (a) The number of composites in the time series representing pixel availability. The maximum number of composites in the time series is 262 for the period between March 2000 and December 2021. (b) Mean number of daily observations within a month used to create the monthly composites as a proxy for uncertainty. The maximum daily observations in a composite are 60 (twice a day every day for a month).

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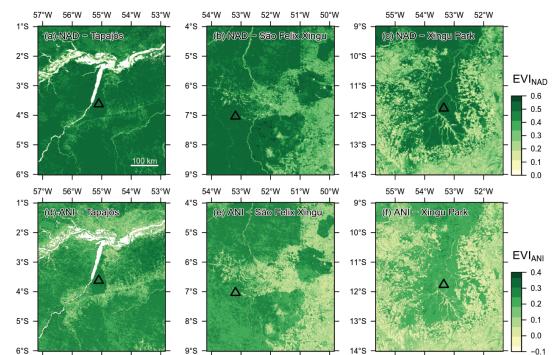
# 309 3. Spatial and temporal distribution of NAD and ANI data across the Amazon forests

310 We selected three experimental areas at the Brazilian Amazon rainforests to show the spatial and 311 temporal distribution of NAD and ANI data (rectangles in Figure 1)To demonstrate the spatial 312 and temporal distribution of NAD and ANI data over the Brazilian Amazon rainforests, we 313 selected three experimental areas (rectangles in Figure 1). These areas show old-growth 314 rainforests with distinct canopy structure and aboveground biomass (AGB) stocks. The AGB increases from semideciduous forests at the Xingu Park (190  $\pm$  19 Mg ha<sup>-1</sup>) and open 315 ombrophilous forests with lianas at the São Felix do Xingu ( $241 \pm 31$  Mg ha<sup>-1</sup>) to dense 316 317 ombrophilous forests at the Tapajós National Forest (288  $\pm$  38 Mg ha<sup>-1</sup>), as estimated by the ESA/CCI AGB map from 2017 (Santoro & Cartus, 2021). These are large-scale AGB estimates 318 319 and may underestimate the true AGB at higher values such as in the Tapajós site. These three 320 sites are also expected to show different phenological dynamics because their selected pixels 321 cover distinct phenoregions in the study reported by Xu et al. (2015).

322 When compared to the nadir-normalized EVI (EVI<sub>NAD</sub>) images (Figures 3a, b, c), the anisotropy 323 EVI (EVI<sub>ANI</sub>) data showed different spatial patterns across sites (Figures 3d, e, f). While the 324 forests over the three sites showed approximately similar EVI<sub>NAD</sub> values (EVI<sub>NAD</sub>  $\approx$  0.50) (Figures 325 3a,b,c), they showed more variability in EVI<sub>ANI</sub> between the Xingu Park (EVI<sub>ANI</sub> > 0.20), São

 $326 \qquad \mbox{Felix do Xingu (EVI_{ANI} > 0.24), and Tapajós (EVI_{ANI} > 0.27) sites (Figures 3d,e,f). This increase and the set of the set of$ 

327 in EVI<sub>ANI</sub> between sites goes into the same direction of the AGB gradient observed from the 328 Xingu Park to the Tapajós National Forest. This result may indicate different forest canopy 329 structures that were not captured in the EVI<sub>NAD</sub> observations, but were captured by the EVI<sub>ANI</sub>. Overall, the EVIANI is high over forests (0.20 to 0.30) and low over pastures and crops (less than 330 331 (0.10). This means large anisotropy between the reflected energy in backward and forward 332 scattering MODIS directions due to the structural complexity of forest canopies. The association 333 between anisotropy and forest canopy structure has been previously shown for the same region in 334 a previous work (Moura et al., 2016).



335

Figure 3 – The spatial distribution in August 2020 (dry season) of the nadir-normalized Enhanced Vegetation Index ( $EVI_{NAD}$ ) is shown in (a), (b), and (c) for the Tapajós National Forest, São Felix do Xingu and Xingu Park, respectively. Corresponding results for the anisotropy EVI ( $EVI_{ANI}$ ) are shown in (d), (e), and (f), respectively. The triangles plotted over (a, b, and c) indicate the sites used to obtain the profiles of Figure 4.

341 From the comparison of different sites (triangles in Figure 3a), we observed that the mean  $EVI_{NAD}$ 342 signal over the time period did not vary much between the selected forests, while the EVIANI 343 varied greatly (Figure 4): Tapajós (mean EVI<sub>NAD</sub> = 0.49, mean EVI<sub>ANI</sub> = 0.27), São Felix do Xingu 344 (mean  $EVI_{NAD} = 0.51$ , mean  $EVI_{ANI} = 0.24$ ), and Xingu Park (mean  $EVI_{NAD} = 0.51$ , mean  $EVI_{ANI}$ 345 = 0.22). Moreover, EVI<sub>NAD</sub> and EVI<sub>ANI</sub> values were moderately positively correlated at Tapajós 346 (r = +0.37), weakly correlated at São Felix do Xingu (r = +0.06), and moderately negatively 347 correlated at the Xingu Park (r = -0.28). The EVI<sub>NAD</sub> and EVI<sub>ANI</sub> are seasonal variability and phase 348 correlation changes from site to site, suggesting that different canopy dynamics processes are 349 likely being captured by the two metrics at the three sites. Understanding exactly what those 350 effects mean for these forests is beyond the scope of this paper. However, it indicates open venues 351 for studying forest functioning using these products. For example, previous studies have shown 352 that EVI<sub>NAD</sub> metrics captured different compositions of leaf ages in the canopies of the central 353 Amazon (Gonçalves et al., 2020).

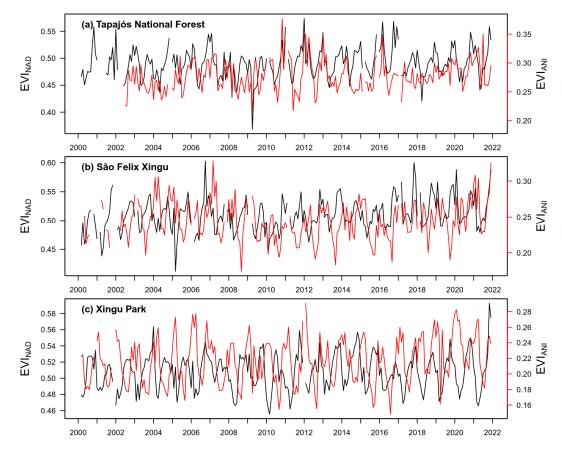


Figure 4 – Time series of AnisoVeg's MODIS Enhanced Vegetation Index (EVI) from 2000 to 2021 for old-growth forests of the (a) Tapajós National Forest; (b) São Felix do Xingu; (c) Xingu Park. The black line indicates the nadir-normalized signal (NAD layer), while the red line represents the EVI anisotropy (ANI layer). The profiles are the mean value of 3 x 3 pixels whose locations are indicated by triangles in Figure 3.

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360 To demonstrate the potential of AnisoVeg for large-scale forest structure inference, we compared 361 the NAD and ANI data against forest height measurements from the Global Ecosystem Dynamics Investigation (GEDI) LiDAR sensor. We found that EVIANI was able to explain up to 55% of 362 363 height variability of Amazon forests according to a simple linear relationship ( $R^2 = 0.55$ , p < 0.01, 364 Figure 5). This is a very strong predicting power for a single variable, considering a simple linear 365 model, especially for satellite passive optical data which are often underrated for forest structure 366 estimates in comparison to Synthetic Aperture Radar (SAR) data. EVINAD was significantly but weakly associated to height variability ( $R^2 = 0.16$ , p < 0.01), reinforcing the increase in 367 368 explanation power owed to the anisotropy metrics built from multi-angle observations. The height 369 data was derived from the GEDI LiDAR sensor aboard the International Space Station. They were 370 obtained more specifically from the product GEDI L2A elevation and height metrics data version 371 2 (footprint size 25 m), acquired from April 2019 to October 2020 (available dates at the time of 372 download). GEDI data were downloaded from Earth Data cloud service system 373 (https://earthdata.nasa.gov). We selected the Relative Height metric at 98th percentile (RH98), 374 which represents the top canopy height. The selected RH98 metric was averaged over each 1-km 375 grid cell, and filtered using a threshold of greater than or equal to 50 shots per  $km^2$  to have a high 376 confidence of reliable height estimation representing the 1-km mean. The AnisoVeg data used for 377 this comparison were based on the same time period as GEDI, and filtered for EVI<sub>NAD</sub> larger than 378 0.35 to exclude non-forested areas. While we only showed the plot for the strongest EVIANI:GEDI 379 relationship in June 2019 (Figure 5), the other months also showed significant (p < 0.01) and

strong relationships with  $R^2$  ranging from 0.36 to 0.55 (mean  $R^2 = 0.46$ ). Future studies should explore relationships using ANI from different months and other indices, alone or in combination with each other, to further understand their significance on-for explaining forest structure. This is important to determine how the anisotropy data can contribute for aboveground biomass and carbon estimates in conjunction with other sources of data such as those from SAR sensors.

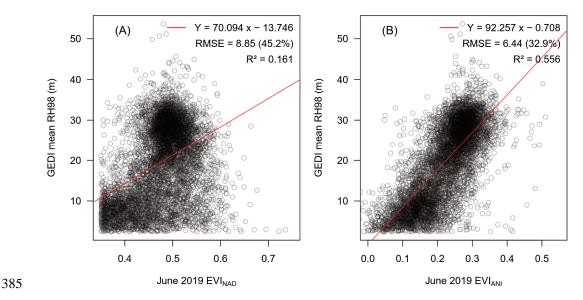


Figure 5 – Relationship between forest height (GEDI mean RH98) and two AnisoVeg layers obtained in June 2019 over the Amazon: (a)  $EVI_{NAD}$  and (b)  $EVI_{ANI}$ . The RH98 metric consists inis the relative height at the 98<sup>th</sup> percentile, which represents the top of canopy height. 7,000 random matching pixels were used in this analysis (1% of 700,000 total matching pixels available), resulting from the filtering of both GEDI and AnisoVeg data. The red line indicates the fitted line by a simple linear model.

392 Terrain illumination is a factor of spectral variability, which can affect EVI<sub>NAD</sub> determination and 393 its relationship with biophysical attributes of vegetation, as shown by previous literature (Huang 394 et al., 2010; Chen and Cao, 2012). Even at 1-km spatial resolution, EVI<sub>ANI</sub> results of Figures 3, 4 395 and 5 can be affected to some extent by terrain illumination effects observed locally at some sites. 396 For instance, topographic effects on EVIANI occurred probably at the São Felix do Xingu site 397 where topographic roughness, observed in SRTM data (results not shown), was coincident with 398 increased EVI<sub>ANI</sub> values in Figure 3E. Furthermore, even in relatively flat terrains, variations in 399 topographic aspect (surface orientation to Sun) can affect the EVI variability in MODIS data 400 because of the different amounts of energy reflected in the NIR towards the sensor by inclined 401 surfaces in the forward and backscattering view directions. Such effects have been observed in 402 southern Brazil with MODIS at 250-m spatial resolution and increased in magnitude at higher 403 spatial resolution data obtained by other sensors (Galvão et al., 2016). Therefore, it may prove 404 useful to include topographic variables in modelling exercises to offset these effects.

405 In a prospective analysis, we also explored the behavior of the two EVI AnisoVeg metrics over 406 the Amazonian phenoregions mapped by Xu et al. (2015). The EVI<sub>NAD</sub> and EVI<sub>ANI</sub> monthly means 407 over different phenoregions highlighted the strong heterogeneity of the Amazonian forests 408 (Figure 6). For instance, the profiles showed strong differences between both metrics from 409 January to September in a phenoregion with well-defined dry and wet seasons (phenoregion one 410 in Figure 6a at the Xingu Park). Large differences between EVI<sub>NAD</sub> and EVI<sub>ANI</sub> were also observed 411 in some phenoregions without a very long dry season in the northwest Amazon (phenoregion five 412 in Figure 6e). On the other hand, EVI<sub>NAD</sub> and EVI<sub>ANI</sub> showed temporal decoupling in phenoregion 413 three located at central-east Amazon (Figure 6c). Overall, while the seasonality of  $EVI_{NAD}$  has 414 been investigated by many studies in the past, the seasonality of  $EVI_{ANI}$  is something to be further 415 explored with the support of auxiliary data (e.g., airborne LiDAR and field campaigns). This is 416 important to better understand the differences in seasonal patterns between both AnisoVeg 417 metrics.

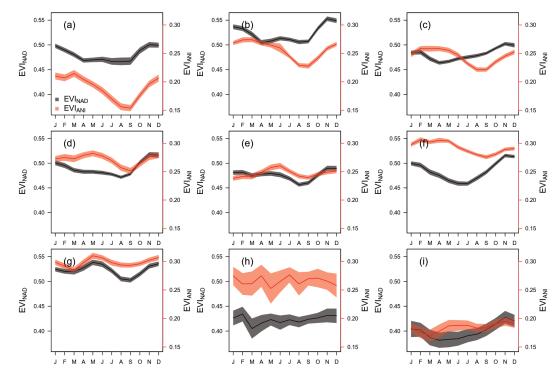


Figure 6 – Monthly means of  $EVI_{NAD}$  (black) and  $EVI_{ANI}$  (red) for nine phenoregions mapped by Xu et al. (2015) in the Amazon. The phenoregions are shown in increasing order from 1 to 9 in corresponding panels (a) to (i). They represent forests with similar seasonality and landscape structure. Solid line and shaded area represent the mean and 95% confidence interval around the mean. The values were extracted from 20 years of data (from 2001 to 2021) for 100 random coordinates within each region, and extracted from 3 x 3 windows of pixels.

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#### 426 **4. Prospective use of the dataset**

427 The NAD layers from the AnisoVeg product have been used in previous studies to explore: the 428 climate drivers of the Amazon forest greening (Wagner et al., 2017); the large-scale Amazon 429 forest sensitivity to drought (Anderson et al., 2018); the structure and dominance of bamboo 430 species in southwest Amazon (Dalagnol et al., 2018); the productivity in a flooded forest in 431 eastern Amazon (Fonseca et al., 2019); the productivity and relationship with Sun-Induced 432 Fluorescence over the Brazilian Caatinga biome (Bontempo et al., 2020); the relationships with 433 leaf-age demography in central Amazon (Gonçalves et al., 2020); and the relationships with fire 434 disturbance and SAR-based Vegetation Optical Depth in southern Amazon (Zhang et al., 2021).

The ANI layers from the AnisoVeg product have been mainly used to characterize Amazon forest structure properties (Moura et al., 2015; 2016). These layers now open new venues of investigation on vegetation, including (but not limited to): the characterization of biophysical attributes of forests, including their seasonality and trends; the assessment of changes in vegetation structure due to natural disturbances or degradation (logging, fire, edge effects); and the evaluation of forest health and productivity (greenness and browning). We expect that this 441 dataset contributes to upscaling studies over large areas of key forest properties such as the AGB 442 and canopy roughness (Foody & Curran, 1994; Saatchi et al., 2008). This information is required 443 for dynamic vegetation models to accurately represent the carbon cycle. This dataset is not limited 444 to study Amazonian forests and can be used to explore other biomes of South America such as 445 the Atlantic Forest, savannas (Cerrado), Caatinga, Chaco, Pantanal, and Pampas. Such studies 446 could improve our understanding of large-scale vegetation functioning, carbon storage, and 447 cycling. Ultimately, they can contribute to refine global ecosystem models, and to obtain accurate 448 estimates of carbon cycle in response to climate and environmental change. Furthermore, 449 auxiliary backward and forward scattering data are also available with the dataset. Beyond the 450 use of the provided ANI layers, this effectively allows the computation of several other multi-451 angular anisotropy indices from the literature (Table 3). The advantage or disadvantage of one 452 specific anisotropy index rather than others is not established in the literature given the range of vegetation applications and the lack of available datasets up to date. We calculated and provided 453 454 only ANI due to its demonstrated relationships with Amazonian forests structure and functioning 455 (Moura et al., 2015; Moura et al., 2016; Hilker et al., 2017). However, we expect other indices, 456 including ratios and normalized differences between the backward and forward scattering 457 components, offer additional possibilities for tropical vegetation studies which should be explored 458 in future studies.Furthermore, auxiliary backward and forward scattering data are also available 459 with the dataset. Beyond the use of the provided ANI layers, this effectively allows the 460 computation of several other multi-angular anisotropy indices from the literature (Table 3), offering the possibility to investigate their use for tropical vegetation studies. 461

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Table 3 – Examples of other multi-angular anisotropy indices that can be further calculated using
layers of the AnisoVeg product. Lambda represents the selected spectral band or vegetation index.
N, HB, and D-F represent nadir-view normalization, hot-spot (backward scattering), and darkspot (forward scattering) estimates, respectively.

Anisotropy Indices	Formula	Reference
Anisotropy index (ANIX)	$\lambda_{HB}$	Sandmeier et
	$\lambda_{DF}$	al. (1998)
Nadir BRDF-adjusted NDVI (NDVI <sub>ISO</sub> )	$NIR_N - RED_N$	Schaaf et al.
-	$NIR_N + RED_N$	(2002)
Hot-spot dark-spot index (HDS <sub>RED</sub> )	$RED_{HB} - RED_{DF}$	Lacaze et al.
	$RED_{DF}$	(2002)
Normalized difference between hot-spot	$NIR_{HB} - NIR_{DF}$	Chen et al.
and dark-spot index (NDHD <sub>NIR</sub> )	$NIR_{HB} + NIR_{DF}$	(2005)
Hot-spot dark-spot NDVI (NDVI <sub>HD</sub> )	$NIR_{HB} - RED_{DF}$	Pocewicz et
	$NIR_{HB} + RED_{DF}$	al. (2007)
Hot-spot-incorporated NDVI (NDVI <sub>HS</sub> )	$NDVI_N \times (1 - RED_{HR})$	Pocewicz et
		al. (2007)
Anisotropy difference (ANI)*	$\lambda_{\mu R} - \lambda_{DE}$	Moura et al.
••		(2015)
Vegetation Structure Index (VSI)	$NDVI_{DF} - NDVI_{HB}$	Sharma et al
	$1 - NIR_{DF}$	(2021)

\*ANI is included in the AnisoVeg product. Source: Adapted from Sharma et al. (2021).

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#### 469 **5. Code and data availability**

All code is available at GitHub (https://github.com/ricds/maiac\_processing) (Dalagnol &
Wagner, 2022). The full dataset can be found at the official AnisoVeg repository at Zenodo

472 (https://doi.org/10.5281/zenodo.3878879) (Dalagnol et al., 2022). The dataset was organized in 473 compressed files (".zip" format) sub-divided by years (currently 2000-2021) and layers (bands 1-474 8, NDVI, and EVI) for both nadir-normalization (code = NAD) and anisotropy (code = ANI). The 475 number of samples layers (code = NO SAMPLES) are also provided. Inside each compressed 476 file there will be 12 image files (".tif' format), one per month, except for the year 2000 which 477 starts in March. The storage size for the whole dataset is 162.6 GbGB. The data have a scale factor 478 of 10,000 to reduce file storage size. Thus, to obtain surface reflectance values of bands or correct 479 range of values for indices, you should divide the layers by 10,000. The exception is the number 480 of samples, which already shows the correct range of values from 0 to 60 observations. The dataset 481 is planned to be updated on a yearly-basis. Auxiliary data that allow the calculation of other 482 anisotropy metrics (listed in Table 3) are included in two separate Zenodo repositories for 483 backward (https://doi.org/10.5281/zenodo.6040300) (Dalagnol, 2022a) and forward scattering 484 (https://doi.org/10.5281/10.5281/zenodo.6048785) (Dalagnol, 2022b), including the selected 485 layers Red, NIR, NDVI and EVI. The  $EVI_{ANI}$  and  $EVI_{NAD}$  layers were also uploaded to the GEE 486 platform using the geeup tool v0.5.3 (Roy, 2022). They can be accessed through the GEE ImageCollection "projects/anisoveg/assets/evi anisotropy" 487 assets and 488 "projects/anisoveg/assets/evi\_nadir", found at 489 <https://code.earthengine.google.com/?asset=projects/anisoveg/assets/evi\_anisotropy> and

490 <https://code.earthengine.google.com/?asset=projects/anisoveg/assets/evi\_nadir>.

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## 492 Author contribution

493 R.D. and Y.M. conceived the presented idea. R.D. designed the methodology with contributions 494 from Y.M. on the anisotropy method. R.D. conducted formal analysis and investigation with 495 contributions from L.G., F.W., N.G., and S.S. Y.W. and A.L. provided the original MODIS 496 (MAIAC) data and support for processing it, Y.Y. and S.S. provided the processed GEDI height 497 data and support to analyze it. R.D. and F.W. developed the code to process the MODIS (MAIAC) 498 data into the products. R.D. conducted data curation of the products. L.A. supervised the project. 499 R.D. wrote the original draft with support from L.G., F.W. and Y.M. All authors read, reviewed 500 and approved the final version of the manuscript.

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# 512 **Conflict of Interest**

513 The authors have declared no conflict of interest.

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