A strontium isoscape of inland southeastern Australia

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Abstract

The values and distribution patterns of the strontium (Sr) isotope ratio $^{87}Sr/^{86}Sr$ in Earth surface materials is of use in the geological, environmental and social sciences. Ultimately, the $^{87}Sr/^{86}Sr$ ratio of soil and everything that lives in and on it is

10 inherited from the rock that is its parent material. In Australia, there are few large-scale surveys of $^{87}Sr/^{86}Sr$ available, and here we report on a new, low-density dataset using 112 catchment outlet (floodplain) sediment samples covering 529,000 km^2 of inland southeastern Australia (South Australia, New South Wales, Victoria). The coarse (<2 mm) fraction of bottom sediment samples (depth ~0.6-0.8 m) from the National Geochemical Survey of Australia were milled and fully digested before Sr separation by chromatography and $^{87}Sr/^{86}Sr$ determination by multicollector-inductively coupled plasma-mass spectrometry.

15 The results show a wide range of $^{87}Sr/^{86}Sr$ values from a minimum of 0.7089 to a maximum of 0.7511 (range 0.0422). The median $^{87}Sr/^{86}Sr$ (± robust standard deviation) is 0.7199 (± 0.0112), and the mean (± standard deviation) is 0.7220 (± 0.0106). The spatial patterns of the Sr isoscape observed are described and attributed to various geological sources and processes. Of note are the elevated (radiogenic) values (≥~0.7270; top quartile) contributed by (1) the Palaeozoic sedimentary country rock and (mostly felsic) igneous intrusions of the Lachlan geological region to the east of the study area; (2) the Palaeoproterozoic

20 metamorphic rocks of the central Broken Hill region; both these sources contribute radiogenic material mainly by fluvial processes; and (3) the Proterozoic to Palaeozoic rocks of the Kanmantoo, Adelaide, Gawler and Painter geological regions to the west of the area; these sources contribute radiogenic material mainly by aeolian processes. Regions of low $^{87}Sr/^{86}Sr$ (≤~0.7130; bottom quartile) belong mainly to (1) a few central Murray Basin catchments; (2) some Darling Basin catchments in the northeast; and (3) a few Eromanga geological region-influenced catchments in the northwest of the study area; these

25 sources contribute radiogenic material mainly by fluvial processes.

1 Introduction

Strontium (Sr) isotopes in Earth systems are useful to elucidate geoscientific, environmental and societal issues, including mineralisation (e.g., Zhao et al., 2021), hydrology (e.g., Christensen et al., 2018), food tracing (e.g., Vinciguerra et al., 2015) and historic migrations of people and animals (e.g., Price et al., 2017; Madgwick et al., 2019). Whilst several countries have

30   developed detailed Sr isotopic maps (Sr isoscapes) of their territory, recent examples of which include Italy (Lugli et al., 2022) and Ireland (Snoeck et al., 2020), no such dataset currently exists for Australia. The largest Sr isoscape to-date in this country is for the Cape York Peninsula in northern Queensland (Adams et al., 2019), which covers approximately 300,000 km^2. In this contribution we publish a new Sr isotopic dataset and isoscape for a part of inland southeastern Australia, across three states, covering agricultural land and remote and arid 'outback' deserts. Additional parts of Australia currently are being investigated
35   to expand this isoscape further.

2 Setting

The study area, the Darling-Curnamona-Delamerian (DCD) area, is located in inland southeastern Australia, across parts of the states of South Australia, New South Wales and Victoria (hereafter abbreviated to SA, NSW and Vic, respectively) between 28 and 37 °South and 138 and 146 °East (Figure 1). The DCD area is one of the 'deep-dive' focus areas targeted for
40   geoscientific data acquisition (https://www.ga.gov.au/eftf/projects/darling-curnamona-delamerian) under the Exploring for the Future program (https://www.ga.gov.au/eftf), a major Australian Government research funding initiative.

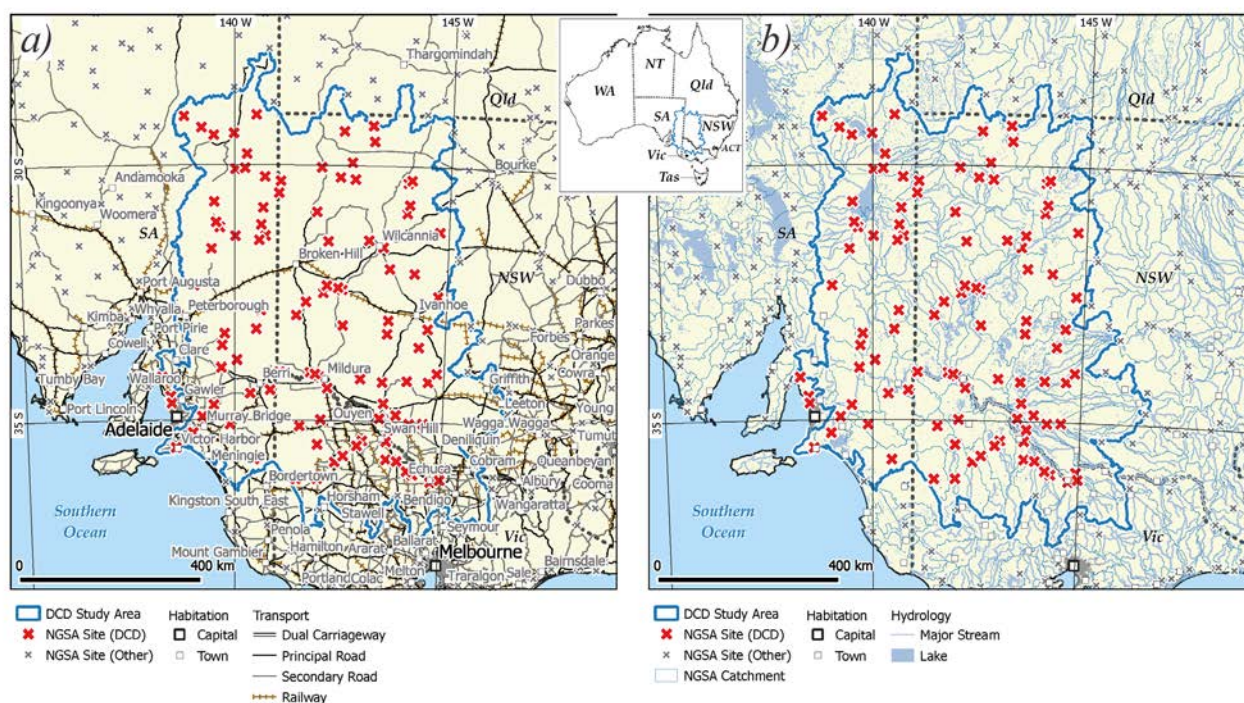



Figure 1: The Darling-Curnamona-Delamerian (DCD) Sr isotope study area (dark blue outline, see inset for location: WA – Western Australia, NT – Northern Territory, Qld – Queensland, NSW – New South Wales, ACT – Australian Capital Territory, Vic –
45   **Victoria, Tas – Tasmania, and SA – South Australia) and National Geochemical Survey of Australia (NGSA) sample locations (crosses) overlain (a) on topographic map showing townships (squares), roads (black or grey), railway (brown), lakes (pale blue), vegetated areas (green) and deserts (red-brown); and (b) on hydrology map showing NGSA catchment boundaries, major streams and lakes. Map projection: Albers equal area.**

The main climate zones in the area are described as 'Warm summer, cold winter' along the southern/southwestern coastal
fringe to 'Hot dry summer, cold winter' over the remainder of the inland area (BOM, 2021a). The vegetation/Köppen zones
are dominated by 'Grassland/Warm (persistently dry)' and 'Desert/Warm (persistently dry)' in the south, and 'Grassland/Hot
(persistently dry)' and 'Desert/Hot (persistently dry)' in the north (BOM, 2021b). The 30-year (1976-2005) average minimum
temperatures and maximum temperatures mostly range 3-18 °C and 18-33 °C (both increasing south to north), respectively
(BOM, 2021c,d). Average annual rainfall over the latest 30-year period available (1986-2015) mostly ranges 50-400 mm/yr
(BOM, 2021e) with precipitation falling mostly in winter along the southern fringe of the area whilst the central and northern
parts are equally arid year-round.

Physiographically the study area includes, clockwise from the southwest, the Adelaide Hills, North St Vincent-Spencer Gulf,
Lower Murray, Lake Frome, Cooper Creek, Bulloo-Bancannia, Warrego, Menindee Lakes, Lachlan, Murrumbidgee, Mid-
Murray, Goulburn-Loddon, Wimmera-Mallee and Millicent Coast physiographic regions (Pain et al., 2011). (Locations of
geographical entities named in this paper are given in the Supplementary Material.) The Murray River is Australia's longest
river, with its source in the Australian Alps at 1430 m above sea level (asl) (on the border between NSW and Vic) being
approximately 2500 km upstream of its mouth near Victor Harbour (SA) where it enters the Southern Ocean. The combined
catchments of the Murray River and that of its northern tributary, the Darling River (the country's third longest river), define
the Murray-Darling Basin, a vast (~1 million km^2), low-relief Cenozoic depression formed in dynamic topographic response
to continental plate motion (Schellart and Spakman, 2015). The Murray-Darling Basin is one of Australia's most productive
agricultural areas where water management, critical to support communities and the environment across its mega-catchment,
is overseen by the Murray-Darling Basin Authority (MDBA, 2022). Topographic elevation (GA, 2008) in the DCD area ranges
from sea level, along the Great Australian Bight region of the Southern Ocean to the southwest, to 1206 m asl, with the highest
elevations (>900 m asl) observed both to the southeast, in the Victorian highlands, and to the northwest, in the Flinders Ranges
of SA; the mean altitude is 139 m asl.

The study area comprises a wide range of geological regions (Blake and Kilgour, 1998; Figure 2a), the main ones ranging from
the Palaeoproterozoic Broken Hill region and Palaeoproterozoic to Mesoproterozoic Mount Painter region, through the
Palaeoproterozoic to Devonian Wonominta and Tibooburra regions, the Neoproterozoic to Cambrian Stuart and Adelaide
regions, the Neoproterozoic to Ordovician Caloola region, the Cambrian Arrowie, Cambrian to Ordovician Kanmantoo and
Cambrian to Devonian Bancannia regions, the Ordovician to Early Carboniferous Lachlan region, the Devonian to
Carboniferous Darling region, the Jurassic to Tertiary Eromanga regions, through to the Late Mesozoic to Quaternary Murray
region. The most dominant surface lithological type is undoubtedly regolith (Eggleton, 2001), the commonly loose (though in
places highly indurated), transported or in-situ weathering product of primary bedrock. Regolith covers over 85% of Australia
(Wilford, 2012) due to relative tectonic stability, low relief, lack of Quaternary glaciation, low rainfall and resistant rocks (Pain
et al., 2012). Comparatively minor areas of outcrop dot and surround the study area, most notably metamorphic rocks in the
central Broken Hill region, sedimentary and felsic to intermediate igneous rocks in the Lachlan region to the east and south,
felsic to mafic igneous rocks and metamorphic rocks in the Mount Painter region to the west, siliciclastic sediments in the

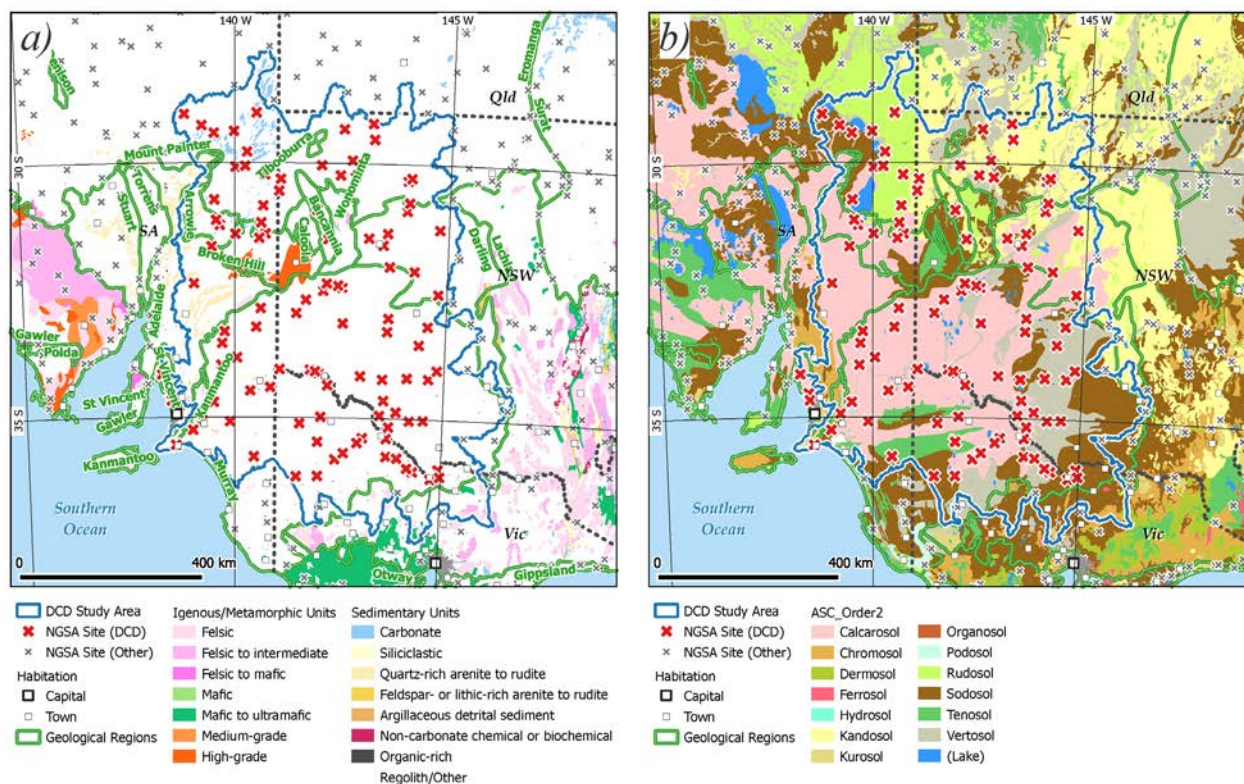
Adelaide and Arrowie regions to the east, and carbonate sediments in the Eromanga region to the north. Very few mafic or ultramafic rocks crop out in the study area, with the main occurrences belonging to the Lachlan region in the southeast.

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Figure 2: The Darling-Curnamona-Delamerian (DCD) Sr isotope study area (dark blue outline), National Geochemical Survey of Australia (NGSA) sample locations (crosses) and geological regions (double green outlines) from Blake and Kilgour (1998) overlain (a) on surface geology from Raymond et al. (2012); and (b) on Australian Soil Classification soil orders from Isbell et al. (2021). Map projection: Albers equal area.

90 The soil types (Orders) encountered in the study area are most commonly calcarosol and vertosol (at 41% and 33% of the sample sites, respectively), followed by sodosol, kandosol and rudosol (8%, 7%, and 5%), and minor chromosol, podosol, and tenosol (~3% to <1%), using the Australian Soil Classification scheme (Isbell et al., 2021; ASRIS, 2021a) (Figure 2b). The major river basins that divide the area mostly belong, clockwise from the southwest, to the Murray-Darling, Lake Eyre, and Bulloo-Bancannia drainage divisions, and (in greater detail) to the Lower Murray, Lake Frome, Bulloo-Bancannia, Warrego,

95 Menindee Lakes, Lachlan, Murrumbidgee, Mid-Murray, Goulburn-Loddon, and Wimmera-Mallee water regions (GA, 1997). Landuse over the area is overwhelmingly grazing native vegetation, with minor nature conservation, grazing modified pastures, and dryland cropping (ASRIS, 2021b), the latter two occurring exclusively in the southern part of the study area.

When mapping and interpreting $^{87}Sr/^{86}Sr$ ratios, it is important to bear in mind the likely processes to may affect them. Pedogenic processes can offset the $^{87}Sr/^{86}Sr$ of soil compared to that of its parent material, as a result of differential dissolution

100 of minerals such as plagioclase and K-feldspar (Blum et al., 1993; Bullen et al., 1997). Here, sediments were collected between

0.6 and 0.8 m below the surface in an attempt to minimise the effect of pedogenic processes on the $^{87}Sr/^{86}Sr$ of sediments post-deposition, however this effect cannot be excluded. In the same way, it is possible that there is an offset between soils in the sediment source region and the parent rocks they derive from. Dust deposition can also affect the soil Sr budget and shift the $^{87}Sr/^{86}Sr$ composition of soil compared to that of its parent material (Stewart et al., 1998; Chadwick et al., 2009). Previous

105 work on soils in the study area has shown that in some cases atmospheric deposition contributes significantly to the soil Sr budget (Green et al., 2004; Mee et al., 2004); in this case, the Sr isotope composition of the soil is not a direct reflection of that of the parent material. However, in other regions of the study area (e.g., Clare Valley), atmospheric deposition plays a much smaller role (Bestland et al., 2003).

3 Material and Methods

110 ### 3.1 Material

This study utilises archive 'catchment outlet sediment' samples collected during the National Geochemical Survey of Australia (NGSA), which covered ~80% of Australia (Caritat and Cooper, 2011, 2016). The sampling philosophy of the NGSA was to collect naturally mixed and fine-grained fluvial/alluvial sediments from large catchments, thereby obtaining representative averages of the main soil and rock types contributing sediment through weathering. This allowed the ultralow sampling density

115 (~1 sample per catchment, which on average is 5200 km^2) to remain representative of large-scale natural variations (Caritat and Cooper, 2011). Catchment outlet sediments are similar to floodplain sediments in the sense that they are deposited during receding floodwaters outside the riverbanks, but with the added complexity that, in Australia, many areas can also experience addition (or loss) of material through aeolian processes. The sampled floodplain geomorphological entities are typically vegetated and biologically active (shrubs, grasses, worms, ants, etc.) thereby making the collected materials true soils, albeit

120 soils developed on transported alluvium parent material.

The NGSA collected samples at two depths, a 'top outlet sediment' (TOS) from a shallow (0.1 m) soil pit approximately 0.8 m x 0.8 m in area, and a 'bottom outlet sediment' (BOS) from a minimum of three auger holes generally drilled within ~10 m of the TOS pit. The auger holes were drilled to refusal or to maximum depth of ~1 m, and the BOS sample was collected, on average, from a top depth of 0.6 m to bottom depth 0.8 m from all augered holes. A field manual (Lech et al., 2007) was

125 compiled to record all sample collection method details, including the site selection process. Sampling for the NGSA took place between July 2007 and November 2009, and the field data were recorded in Cooper et al. (2010).

In the laboratory, the samples were air dried at 40 °C for a minimum of 48 h (or to constant mass), homogenised and reduced by riffle splitting, sieved to two different grain-size fractions (a coarse fraction of <2 mm and a fine fraction of <75 µm), and further prepared (see Caritat et al, 2009) for the comprehensive geochemical analysis program of the NGSA (see Caritat et al.,

130 2010). For Sr isotope analysis, an aliquot of minimum ~1 g of sample was milled to a fine powder using a carbon steel ring mill or, for a few samples only, an agate micromill. The main sample type selected for the present Sr isotope study was BOS

<2 mm to be as representative as possible of the geogenic background unaffected by modern landuse practices and inputs (e.g., fertilizers).

Overall 112 NGSA BOS <2 mm samples were analysed for Sr isotopes as detailed in the Methods Subsection below. Those

135 samples originate from within 102 NGSA catchments, which together cover 529,000 km^2 of inland southeastern Australia (see Figure 1).

3.2 Methods

Samples were prepared and analysed for Sr isotopes (^{87}Sr/^{86}Sr) at the Wollongong Isotope Geochronology Laboratory (WIGL). Approximately 50 mg of sample was weighed and digested in a 2:1 mixture of hydrofluoric and nitric acids. All reagents used

140 were Seastar Baseline® grade with Sr concentrations typically <10 parts per trillion. Following digestion, samples were re-dissolved in aqua regia (twice if needed) in order to eliminate any fluorides, followed by nitric acid twice. Finally, samples were re-dissolved in 2M nitric acid prior to ion exchange chromatography. Strontium was isolated from the sample matrix using automated, low-pressure chromatographic system Elemental Scientific prepFAST-MC™ and a 1 mL Sr-Ca column (Eichrom™) (Romaniello et al., 2015). The Sr elutions were re-dissolved in 0.3M HNO$_3$. Strontium isotope analysis was

145 performed on a Thermo Scientific Neptune Plus multicollector-inductively coupled plasma-mass spectrometry (MC-ICP-MS) at WIGL. The sample introduction system consists of an ESI Apex-ST PFA MicroFlow nebulizer with an uptake rate of ~0.1 mL/min, an SSI Quartz dual cyclonic spray chamber, jet sample and X skimmer cones. Measurements were performed in low-resolution mode. The instrument was tuned at the start of each session with a 20 parts per billion Sr solution, and sensitivity for ^{88}Sr was typically around 4 V. Masses 88, 87, 86, 85, 84 and 83 were collected on Faraday cups. Instrumental mass bias

150 was internally corrected using measured ^{88}Sr/^{86}Sr. Masses 85 and 83 were used to correct for the isobaric interference of ^{87}Rb and ^{86}Kr, respectively.

Maps were prepared using the open software QGis (version 3.16.14-Hannover) and applying an Albers equal area projection. Symbology for displaying ^{87}Sr/^{86}Sr data was either point data classified by eight quantiles (12.5% of the data each; green = low to high = red) at the sampling site, or attributing this same value and colour to the whole catchment from which the outlet

155 sediment comes, reflecting the sampling medium, catchment outlet sediment, being a representative sample of the average materials in the catchment (see Subsection 3.1).

3.3 Quality Assessment

NIST SRM987 was used as a secondary standard and measured after every five samples to assess accuracy during analysis. Accuracy of the whole procedure was assessed by processing USGS reference material BCR-2 (Plumlee, 1998). The mean ±

160 2se ^{87}Sr/^{86}Sr for BCR-2 in this study is 0.704961 ± 35 (n = 13), within error of the value in Jweda et al. (2016) (0.704500 ± 11). Total procedure blanks ranged between 0.025 and 0.245 ng Sr (n = 12).

Ten field duplicate sample pairs (collected at a median distance of ~80 m from one another on the same landscape unit, see Lech et al., 2007) were analysed for ^{87}Sr/^{86}Sr in the BOS <2 mm sample and returned a median relative standard deviation of

0.10%. The relative standard deviation from field duplicates includes natural variability (mineralogical/chemical heterogeneity

165   of the alluvial deposit), as well as sample collection, preparation and analysis uncertainties.

Overall, we feel that the quality of the $^{87}Sr/^{86}Sr$ data presented herein is adequate for the purpose of regional mapping and that reporting $^{87}Sr/^{86}Sr$ data to the third decimal place with an indicative fourth decimal place is appropriate for this work. This relatively low precision is attributed to heterogeneity of the alluvial deposits, since precision relating to sample preparation and analysis for Sr isotopes is at the fifth decimal place (see results for BCR-2 above).

170   4 Results and Discussion

The DCD region $^{87}Sr/^{86}Sr$ results are summarised in Table 1 and the full dataset (n = 112) is available from https://portal.ga.gov.au/xxxxxx. The observed $^{87}Sr/^{86}Sr$ values range from a minimum of 0.7089 to a maximum of 0.7511 (range 0.0422). The median $^{87}Sr/^{86}Sr$ (± robust standard deviation) result is 0.7199 (± 0.0112), and the mean (± standard deviation) is 0.7220 (± 0.0106). Kurtosis and skewness are 0.28576 and 1.0285, respectively. The whole-dataset distribution

175   characteristics indicate a bimodal platykurtic (i.e., displaying few outliers) distribution with a heavy positive (high) tail (Figure 3). Indeed, the Tukey boxplot (Tukey, 1977), shows only one structural outlier at the maximum value. The modes of the two main populations are at ~0.712 and ~0.721.

Table 1: Summary statistics (count, minimum, 25th percentile, median, robust median, mean, standard deviation, 75th percentile,
180   **maximum, range, variance, kurtosis, skewness) of the $^{87}Sr/^{86}Sr$ data from the Darling-Curnamona-Delamerian (DCD) Sr isotope study area, southeastern Australia.**

n	Min	25%	Med	Rob SD	Mean	SD	75%	Max	Range	Var	Kurt	Skew
112	0.7089	0.7130	0.7199	0.0112	0.7220	0.0106	0.7274	0.7511	0.0422	1.11E-4	0.2856	1.0285

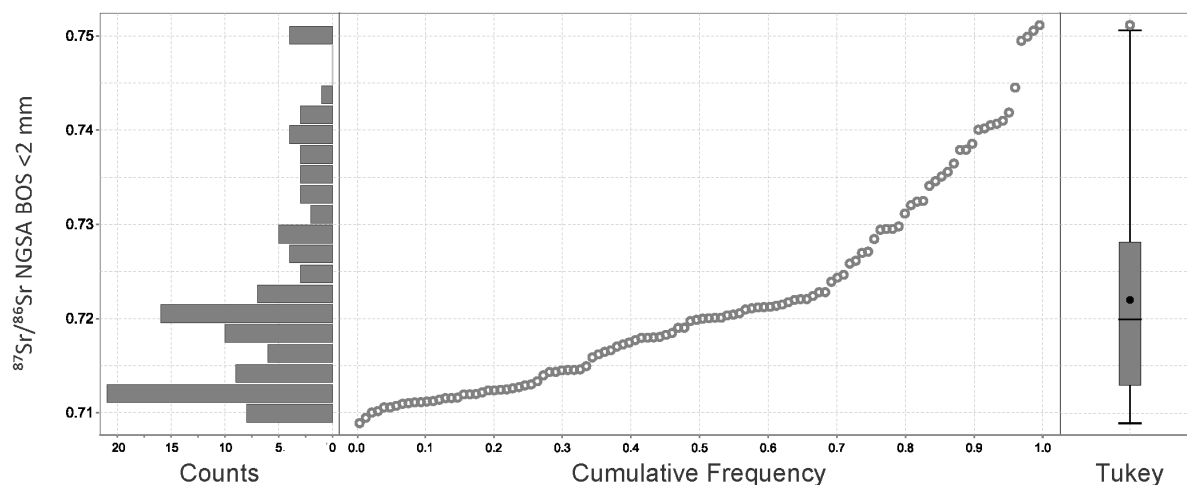



Figure 3: Univariate distribution of the ^{87}Sr/^{86}Sr data from the Darling-Curnamona-Delamerian (DCD) Sr isotope study area: (a)
185 **histogram (20 bins 0.002 wide); (b) cumulative frequency plot; and (c) Tukey boxplot (mean = black dot).**

The observed ^{87}Sr/^{86}Sr values range from close to modern seawater ^{87}Sr/^{86}Sr (0.709), also reflecting modern carbonate rock
and mudstone, to values characteristic of older, Rb-rich felsic igneous rocks, such as granite and rhyolite; the median and mean
values are typical of monzogranite to quartzite main rock types (Bataille and Bowen, 2012, figure 5). In the following
Subsections we discuss potential sources for the more elevated ^{87}Sr/^{86}Sr values in the area, followed by an analysis of potential
190 depleted sources, the variation of ^{87}Sr/^{86}Sr along the Murray River corridor, the variation of ^{87}Sr/^{86}Sr in different geological
regions within the DCD area, and finally compare the ^{87}Sr/^{86}Sr data to existing geochemical and mineralogical data for the
same NGSA sites.

4.1 Sr isoscape mapping

From the ^{87}Sr/^{86}Sr value reported at any given sampling location, we can infer a representative catchment ^{87}Sr/^{86}Sr composition
195 for the drainage area directly upstream of it (Figure 4) because sediments sampled at this location (outlet of cathment) represent
an average of sediments eroded upstream and transported to the sampling location (Ottesen et al., 1989; Bølviken et al., 2004).
This approach allows Sr isoscape mapping of a large area despite a low sampling resolution. However, in addition ot the
limitations relating to weathering and aeolian inputs mentioned above, one needs to bear in mind that this method assumes that
all areas of the landscape draining into the sampling location contribute equally to the sediment budget, which may not be
200 accurate. These limitations need to be taken into consideration before taking the reported ^{87}Sr/^{86}Sr at face value as proxies for
the Sr isotope composition of the basement in the mapped region. If using this dataset to study past or present animal mobility,
another limitation is that we measured the isotopic composition of bulk Sr in sediments, not the bioavailable Sr. Nevertheless,
the dataset presented here can still be used as a guide, or indeed an input to machine learning models, for geological,
archaeological, palaeontological and ecological studies.

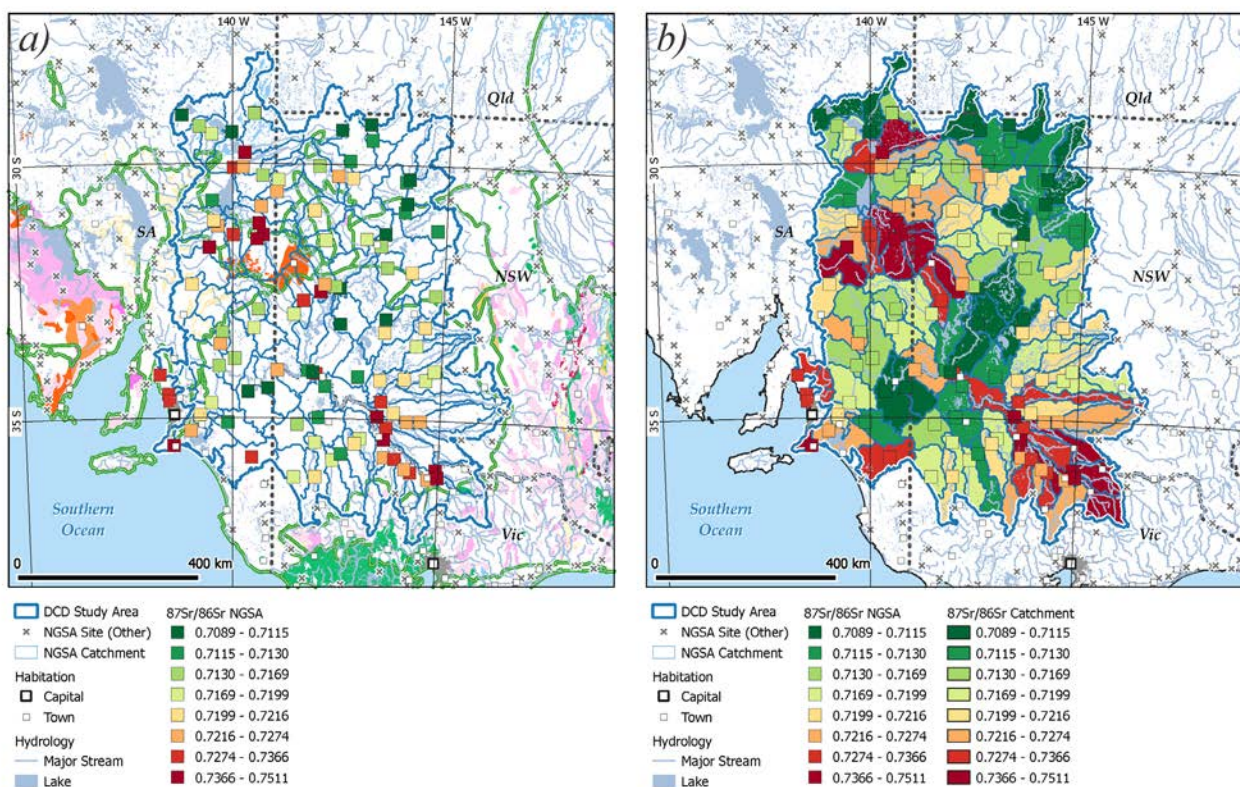



Figure 4: Strontium isoscape for the Darling-Curnamona-Delamerian (DCD) Sr isotope study area with data points classed by quantiles and overlain (a) on geological regions (double green outlines) from Blake and Kilgour (1998) and surface geology from Raymond et al. (2012) (see Fig. 2a for geology legend); and (b) on catchments coloured by same colour ramp. Map projection: Albers equal area.

4.2 Identification of radiogenic ^{87}Sr/^{86}Sr source regions

Spatial patterns observed on the ^{87}Sr/^{86}Sr map (Figure 4) indicate inputs and transport of elevated ^{87}Sr/^{86}Sr signatures down several waterways. The most prominent such 'train' of dispersion is along the Murray River that forms the state border between NSW and Vic in the southeastern part of the study area. Here several ^{87}Sr/^{86}Sr values range from 0.7274 to 0.7511 (i.e., in the top quartile of the dataset), including the maximum value from this survey. On closer inspection, it is evident that the elevated ^{87}Sr/^{86}Sr signatures come down not only the Murray River (three samples), but also from tributaries that join the Murray River and carry sediments down catchments from upland areas both south and north of the border river. Examples of such streams are (from east to west) the Goulburn, Bendigo, and Loddon rivers/creeks located south of the Murray River, and the Wakool, Edward, and Murrumbidgee rivers/creeks to the north of it (see Supplementary Material). Both the Murray and tributary rivers drain abundant felsic igneous rocks hosted in the Ordovician turbidites of the Lachlan Fold Belt to the east and in the Victorian highlands to the south and southeast (both part of the Lachlan geological region), which are likely candidates for sourcing the elevated ^{87}Sr/^{86}Sr minerals transported down catchment. Present-day ^{87}Sr/^{86}Sr ratios for those sediments and intrusions have

indeed been reported around and over 0.72 (e.g., Keay et al., 1997; Chappell et al., 2000; Hergt et al., 2007). Minor mafic lithologies, e.g. at Dookie east of Shepparton (Broken Creek catchment) with lower $^{87}Sr/^{86}Sr$ reported values of 0.704 to 0.707 (Foster et al., 2009), are not of sufficient extent to significantly dampen the elevated $^{87}Sr/^{86}Sr$ signature at the catchment outlet.

225   A lone similarly elevated (top-quartile) sample is located in the south central part of the study area and comes from a dune-dominated landscape devoid of principal drainage (Ngarkat Conservation Park in SA). Only a couple of very small areas of Cambrian-Ordovician felsic igneous rocks crop out in the west of this catchment, possibly suggesting that the elevated $^{87}Sr/^{86}Sr$ signal is inherited by aeolian input from source region(s) further afield. Indeed, the strongest average winds over the period 2004-2008 at this site, which tend to occur in winter, are from the northwest (BOM, 2021f). This upwind direction has abundant

230   Adelaidean-age (meta)sediments, felsic and intermediate igneous rocks as well as medium- and high-grade metamorphic rocks ($^{87}Sr/^{86}Sr$ commonly >0.73), e.g., on the Yorke and Eyre peninsulas and in the Gawler geological region, which are potential sources of elevated $^{87}Sr/^{86}Sr$ mineral grains potentially transported as windblown dust (Figure 4; Foden et al., 2001; Bestland and Forbes, 2009).

Along the Southern Ocean coast, both north and south of Adelaide, four top-quartile samples occur along the Inman, Gawler,
235   Light, and Wakefield rivers, which are in more urbanised/industrialised environments compared to the rest of the study area. These samples are likely also influenced by aeolian transport and deposition of radiogenic material from the Yorke, Eyre and Gawler source areas mentioned above.

South and north of the Broken Hill geological region, near the centre of the study area, similarly elevated $^{87}Sr/^{86}Sr$ values (≥0.7274) are observed in several catchments: Pine and Stephens creeks to the south, and Siccus River, and Digby, Jack, and
240   Childs dam channels (the latter three are from low-relief and featureless landscapes with poorly defined natural drainages), and Eurinilla Creek to the north. All of these catchments drain felsic igneous and/or high-grade metamorphic lithologies in the Palaeoproterozoic Broken Hill domain that would likely be the sources of elevated $^{87}Sr/^{86}Sr$ signatures (Pidgeon, 1967; Page and Laing, 1992) in those sediments. The radiogenic signature of the Broken Hill basement rocks can also be recognised in the regional groundwater (Caritat et al., 2005).

245   Two samples north of Lake Frome also have top-quartile $^{87}Sr/^{86}Sr$ signatures. One of them is located at the bottom of the catchment draining Mount Painter in the northeastern Flinders Ranges, an area of felsic igneous and medium-grade metamorphic rocks, which would contain prime candidate lithologies to produce sediment with a radiogenic Sr isotope signature (Elburg et al., 2003; McLaren et al., 2006). The second sample, located near Bakers Bore at the edge of the Strzelecki Desert, in a field of southwest-northeast oriented linear dunes, likely represents the $^{87}Sr/^{86}Sr$ signature of wind-transported
250   material from the same Mount Painter source region and/or the outcropping Proterozoic rocks of the Broken Hill Domain discussed above, broadly consistent with dominant wind directions (BOM, 2021f).

4.3 Identification of unradiogenic $^{87}Sr/^{86}Sr$ source regions

Whereas much of the attention above has been paid to identifying the potential source rocks or regions for the more radiogenic $^{87}Sr/^{86}Sr$ catchment sediment signatures observed in the DCD study, it is also instructive to look at the catchments with lower

255    $^{87}Sr/^{86}Sr$. Catchments with $^{87}Sr/^{86}Sr$ values <~0.7130 (lower quartile) occur in the central Murray Basin away from the main
Murray River. To the south of the river, several low-relief (essentially internally draining) catchments with virtually no bedrock
outcrop have $^{87}Sr/^{86}Sr$ values (from west to east) at 0.7129 and 0.7126 (field duplicates; Karoonda catchment), 0.7112 (Buniyp
Plain catchment), 0.7095 (Berri catchment), and 0.7120 (Callaghan Plain catchment). The Callaghan Plain catchment is abutted
to the south by the Ouyen catchment ($^{87}Sr/^{86}Sr$ = 0.7116) and the Woomelang catchment to its south (0.7122); both of which

260   are also low relief, locally featuring dunes and/or man-made drainage channels rather than natural stream lines. All of these
$^{87}Sr/^{86}Sr$ values are at the lower end of the typical Murray River sediments as discussed in the next Subsection.

To the north of the Murray River, the Great Darling Anabranch and Darling River have signatures of 0.7120, and 0.7116 and
0.7114 (field duplicates), respectively. The Arumpo catchment has a value of 0.7127, whilst to its north the Karoola, Garnpung,
and Talyawalka catchments have $^{87}Sr/^{86}Sr$ values of 0.7114, 0.7110, and 0.7112, respectively. All of these catchments lack

265   mapped rock outcrop and consist therefore almost entirely of Darling Basin sediment, with an unknown potential contribution
from aeolian material in places. This is entirely consistent with the assessment of Gingele and De Deckker (2005) that typical
source rocks in the New England Fold Belt (where Darling River tributaries mainly originate from) are Tertiary basalts with
$^{87}Sr/^{86}Sr$ ratios of 0.703-0.705 and Permian granites with ratios around 0.7127.

In the northeast of the DCD study area, there are several catchments also with low $^{87}Sr/^{86}Sr$ signatures: 0.7102 (Paroo

270   Overflow), 0.7124 (Tambua), 0.7116 (Paroo Overflow 2), 0.7108 (Cuttaburra), 0.7125 (Paroo), and 0.7106 (Cuttaburra
Channels) toward the north-northeast from Wilcannia. In the central north, four further catchments have lower quartile $^{87}Sr/^{86}Sr$
values: Caryapundy (0.7111), Bootra (0.7120), Feeder (0.7125), and Berawinnia (0.7106) catchments.

Finally, in the northwest of the study area where the Eromanga geological region encroaches onto the DCD area, a sand dune-
dominated catchment on the edge of the Strzelecki Desert east of Lake Blanche has low $^{87}Sr/^{86}Sr$ (0.7100), as does the

275   Cooryanna catchment (0.7089) further west.

4.4 Murray River profile

A profile of the catchment sediments' $^{87}Sr/^{86}Sr$ value along the Murray River (and tributaries) is shown in Figure 5 for sampling
sites within a 5 km buffer and a 10 km buffer from the Murray River. The profile indicates a marked decrease in $^{87}Sr/^{86}Sr$ from
the highlands (0.74-0.75) to near the mouth of the river (~0.71), spanning nearly the whole data range for the DCD study area.

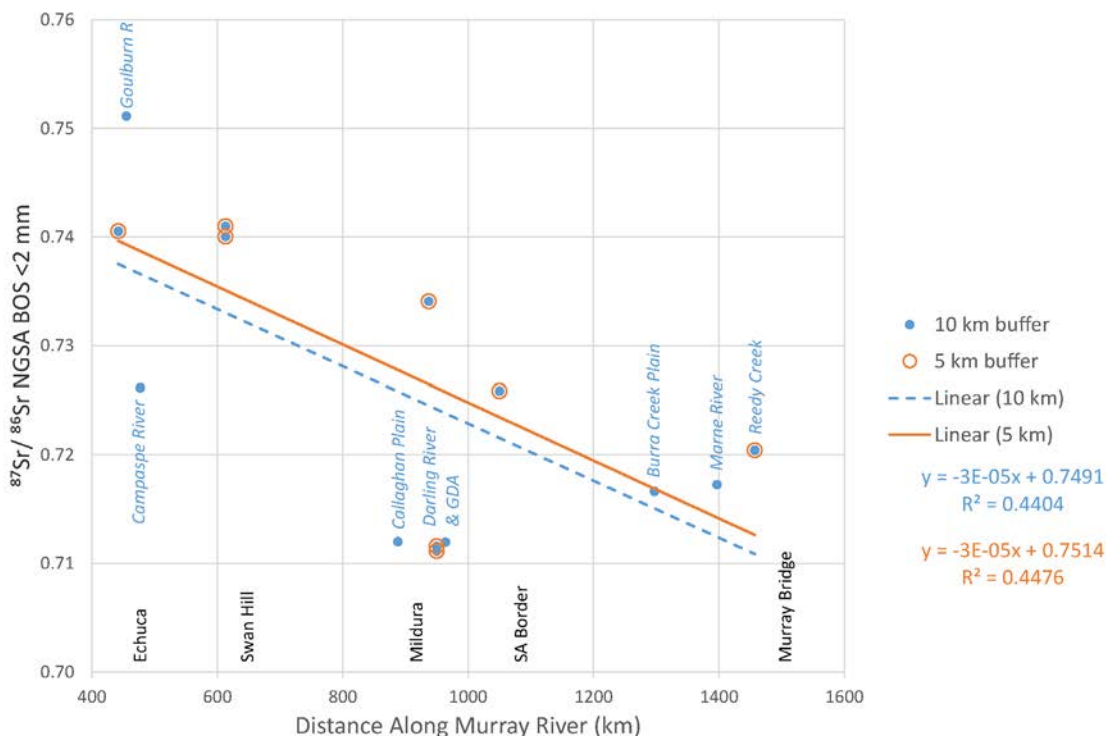
280   The 5 km and the 10 km buffered sampling points give linear regressions with similar correlation coefficients (R = -0.6690
and -0.6636, respectively) and identical slopes (both -3 × 10^{-5}). Both regressions also suggest a primary $^{87}Sr/^{86}Sr$ of source
material at the source of the Murray River (intercept) of 0.749-0.751, which is in agreement with the isotopic signature of
felsic igneous and sedimentary rocks in the Lachlan geological region (e.g., Keay et al., 1997; Chappell et al., 2000; Hergt et
al., 2007).

285

Figure 5: ^{87}Sr/^{86}Sr profile of NGSA catchment sediment sampling sites located within 5 km and 10 km buffers of the Murray River (open and closed circles, respectively) vs downstream distance along the Murray River (east to west). Sites not directly sampling a Murray River catchment are labelled (in italics) with the name of the tributary (R = River; GDA = Great Darling Anabranch) or geomorphic entity. Least squares regression lines for the 5 km and 10 km buffers are shown as solid and dashed lines, respectively,
290 **and equations are given. Township and South Australia (SA) border positions are shown on the X-axis.**

Along this profile, a few tributaries have clearly anomalous ^{87}Sr/^{86}Sr signatures, for instance the Goulburn River and the Campaspe River in the highlands are higher (~0.751) and lower (~0.726), respectively, than the Murray River sediment signature in this part of the Basin (~0.740). Further downstream, the numerous channels and low-relief landscape from the Callaghan Plain contribute low ^{87}Sr/^{86}Sr material to the main Murray valley from the south. Next, the Darling River confluence
295 also contributes less radiogenic sediments (~0.711) from the north.

These findings are consistent with the study of the fluvial suspended particulate matter by Douglas et al. (1995), which found more radiogenic values in clay and coarser (>0.2 µm) fraction material from the Murray River (~0.722 to 0.760) than from the Darling River (~0.709-0.711). Gingele and De Deckker (2005) studied the Sr isotopic composition of fluvial clay-sized sediments (<2 µm) of the Murray and Darling Basins, and also found the former to be more radiogenic and variable (0.726-
300 0.775) than the latter (0.708-0.717). Aeolian dust collected in the Murray and Darling Basins (De Deckker et al., 2014) similarly shows a radiogenic divide between these source regions, with Murray Basin dust spanning a wide range of more radiogenic ^{87}Sr/^{86}Sr (0.716-0.776) compared to the Darling Basin dust (0.708-0.714).

Further downstream, closest to the mouth of the river, sediments are being transported down from the west with a more radiogenic signature (~0.720) relative to the Murray River sample(s) further upstream; this signature is consistent with rocks
305   from the Kanmantoo geological region (Gray, 1990; Bestland and Forbes, 2009; Haines et al., 2009) shedding sediment down Reedy Creek.

4.5 Isotopic signature of the geological regions

In consideration of the above discussion on more radiogenic and less radiogenic source regions for the minerals found in the NGSA sediments of the DCD, Figure 6 shows a clear split between these regions in the $^{87}Sr/^{86}Sr$ vs 1000/Sr variable space.
310   By and large, the sediments provenanced from, or deposited in, the Darling Basin and Eromanga Basin (only the southern part of which is considered in this study) appear to have relatively low $^{87}Sr/^{86}Sr$ ratios (mostly <0.7250), in agreement with previous studies cited above. At the same time, these sediments have a widely variable Sr content (1000/Sr ranging from 1.2 to 33.6 g.µg^{-1}). In contrast, NGSA sediments collected in catchments draining[1] the other main geological regions represented in the DCD area (Adelaide, Arrowie, Broken Hill, Kanmantoo, Lachlan, and Murray geological regions; open symbols in Figure 6)
315   can have more radiogenic $^{87}Sr/^{86}Sr$ signatures (up to 0.750). These sediments also have a more limited Sr content range (1000/Sr ranging mostly from 0.4 to ~21 g.µg^{-1}). Generalised trends for these two populations have been suggested by the qualitative arrows in the scatterplot.

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[1] Samples are attributed a geological region here not based on which region they are necessarily located in (sample site), but based on which region is most represented in their catchment
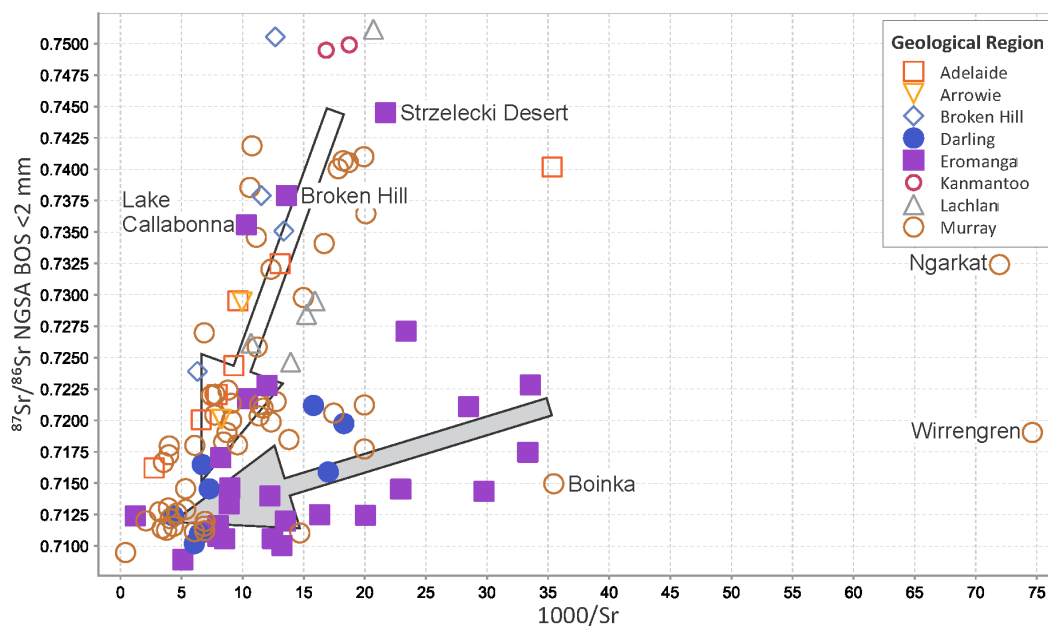
Figure 6: Scatterplot of $^{87}Sr/^{86}Sr$ vs 1000/Sr in the NGSA BOS <2 mm samples from the Darling-Curnamona-Delamerian (DCD) Sr isotope study area, grouped by main geological region each catchment is draining.

Three samples from the Murray geological region, however, have much lower Sr contents (1000/Sr between 35 and 75 g.µg^{-1}) than the others (Figure 6); closer examination of these unusual cases indicates that they are from sandy deserts near the Vic-SA border in the southwestern DCD area (Boinka at the southern edge of Sunset National Park, Ngarkat in the Ngarkat Conservation Park, and Wirrengren in the Big Desert). Here, aeolian input of quartz grains may be or may have been likely active, diluting the geochemical and mineralogical concentrations of Sr-bearing components in the samples. Indeed these three sites contain over 90 wt% SiO_2, which is over the 75th percentile nationally for the NGSA (Caritat and Cooper, 2011) and among the highest silica content of all the DCD samples. Since the aeolian quartz input is likely exogenous (from a different source to the provenance of the silt and clay material in the sample), we can reasonably speculate that the fluvial component of these materials would be significantly richer in Sr and therefore have lower 1000/Sr values.

Three catchment samples draining mainly the Eromanga geological region, conversely, fall within the general Adelaide-Broken Hill-etc. trend (open symbols and arrow in Figure 6). These come from the Broken Hill, Lake Callabonna, and Strzelecki Desert catchments in the central-northern part of the study area. The samples are located either downstream from the Broken Hill highlands or downwind from the Mount Painter and Adelaide regions. They are thus likely influenced by these more radiogenic source regions even though physically located in the Eromanga geological region, which typically has a lower $^{87}Sr/^{86}Sr$ sediment signature.

To refine the differentiation of $^{87}Sr/^{86}Sr$ signature for the various geological regions, boxplots of the data sub-grouped by these regions is presented in Figure 7. In Figure 7a, Tukey boxplots are shown first for the whole dataset (n = 112), then in succession by the drained geological region in order of increasing median (or, in two cases, unique) value: Darling (n = 8), Wonominta

(2), Eromanga (26), Bancannia (1), Murray (54), Arrowie (2), Adelaide (7), St Vincent (1), Lachlan (5), Broken Hill (4), and

340   Kanmantoo (2) geological regions. A histogram indicating the contribution of each geological region to the general population distribution is shown in Figure 7b. With the few exceptions explained above, both diagrams support a general model with two main sources of fluvial and aeolian sediment in the DCD area, a more radiogenic source ($^{87}Sr/^{86}Sr$ between 0.720 and 0.751) consisting generally of the older (Palaeoproterozoic to Early Carboniferous) geological regions (Kanmantoo, Broken Hill, Lachlan, Adelaide and Arrowie) versus a less radiogenic source ($^{87}Sr/^{86}Sr$ between 0.709 and 0.722) consisting of generally

345   younger (Devonian to Tertiary) geological regions (Eromanga and Darling). The most expansive and youngest (Late Mesozoic-Quaternary) Murray geological region represents a transition between these two endmembers, as, depending on sampling location, it is the depositional site for material mainly contributed by either the radiogenic or the unradiogenic endmember source.
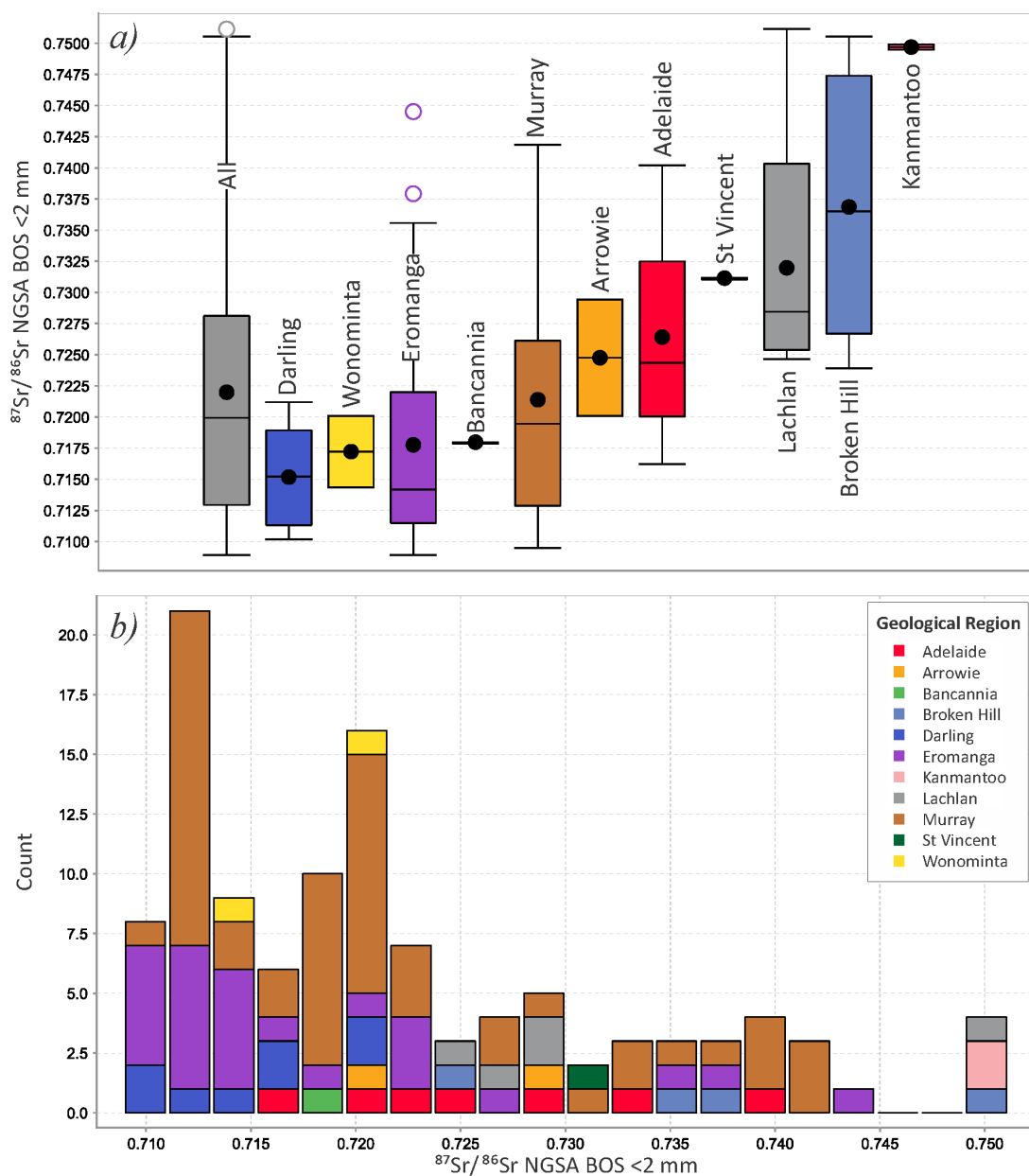
Figure 7: Univariate distribution representations of the $^{87}Sr/^{86}Sr$ data from the Darling-Curnamona-Delamerian (DCD) Sr isotope study area by geological region: (a) Tukey boxplots of, left to right, the whole dataset first (grey), then sorted by increasing median (Darling, Wonominta, Eromanga, Bancannia, Murray, Arrowie, Adelaide, St Vincent, Lachlan, Broken Hill, and Kanmantoo geological regions); (b) histogram by the above subgroups.

The smaller Palaeoproterozoic-Devonian Wonominta region, drained by two of the catchments sampled here, yields $^{87}Sr/^{86}Sr$

values (0.7144 and 0.7201) uncharacteristic of the contemporaneous radiogenic endmember identified above, but intermediate

between it and the Mt Wright metavolcanics of alkali basalt affinity (0.704-0.706; Plimer, 1985; Zhou and Whitford, 2007)

found in the Wonominta block. In the absence of published $^{87}Sr/^{86}Sr$ data for other lithologies, particularly felsic volcanic and (meta)sedimentary lithologies, from the Wonominta block, it is impossible to determine if the two catchment sediment values reported here are typical of a weighted average Wonominta block Sr isotope signature or, instead, represent a mixture of
360   material originating from it and from either of the two endmembers discussed above.

4.6 Comparison to geochemistry and mineralogy

The $^{87}Sr/^{86}Sr$ values in the DCD area discussed here are loosely correlated with the concentrations of major elements/constituents potassium (Spearman correlation coefficient $r_S = 0.35$), calcium (-0.33), and loss on ignition (a proxy for organic matter; -0.32) in the same BOS coarse NGSA sediment samples (Caritat and Cooper, 2011). (Note that for n = 112,
365    $|r_S|$ values > 0.24 have p = 0.001.) A number of trace elements similarly have loose correlations with the $^{87}Sr/^{86}Sr$ values in the DCD area, most notably the total concentrations of thorium (0.48), tungsten (0.44), lanthanum (0.42) and light rare earth elements praseodymium, cerium and neodymium (0.38 to 0.35), rubidium (0.35), lead (0.35), and Sr itself (-0.52) (Caritat and Cooper, 2011). In terms of bulk mineralogy, the $^{87}Sr/^{86}Sr$ values of the DCD samples correlate best with the abundances of plagioclase (0.44), and illite-muscovite (0.43) in the TOS bulk NGSA sediments (Caritat and Troitzsch, 2021). Therefore,
370   weak spatial distribution similarities exist between the $^{87}Sr/^{86}Sr$ maps and the geochemical/ mineralogical maps for the above parameters.

5 Data Availability

The new spatial Sr isotope dataset for the DCD region is publicly available through the Geoscience Australia portal (Caritat et al., 2022; https://dx.doi.org/10.26186/146397).

6 Conclusions

One hundred and twelve catchment outlet sediment samples from the National Geochemical Survey of Australia archive were analysed for strontium (Sr) isotopic composition ($^{87}Sr/^{86}Sr$). Geographically, this study covers 529,000 km^2 of the Darling-Curnamona-Delamerian (DCD) area in inland southeastern Australia, and is dominantly flat and low-lying with minor bedrock outcrop. The bottom outlet sediment (BOS) samples were retrieved mostly by augering overbank or floodplain landforms near
380   the outlet of large catchments to, on average, 0.6 to 0.8 m depth. Total digestion of milled <2 mm grain-size fractions from these sediments yielded a wide range of $^{87}Sr/^{86}Sr$ values from a minimum of 0.7089 to a maximum of 0.7511 (range 0.0422), with a median (± robust standard deviation) of 0.7199 (± 0.0112), and a mean (± standard deviation) of 0.7220 (± 0.0106). A map of the $^{87}Sr/^{86}Sr$ distribution (isoscape) reveals spatial patterns reflecting the sources and transport pathways of the sediment carried down rivers, which drain mainly from the east and the north to the west/southwest, and, to a lesser degree,
385   by wind, which blows dominantly from the west/northwest. Three main source regions of relatively radiogenic ($^{87}Sr/^{86}Sr$

≥~0.7270) clastic material are recognised. First, the Palaeozoic felsic igneous and sedimentary rocks of the Lachlan geological region to the east and southeast of the DCD area contribute fluvial sediment that is transported west by the Murray River and several tributaries. Second, in the central DCD area, alluvium shed by the topographically proud outcrop of Palaeoproterozoic metamorphic rocks of the Broken Hill geological region disperses radially to the north, west, and southeast. Third, radiogenic

390   material sourced from the Proterozoic to Palaeozoic rocks of the Kanmantoo, Adelaide, Gawler and Painter geological regions to the west of the area disperses mostly by aeolian processes toward the east and southeast. These radiogenic sources contrast with the relatively unradiogenic ($^{87}Sr/^{86}Sr$ ≤~0.7130) sediments found in a few, mostly internally draining, catchments in the central Murray Basin, some Darling Basin catchments in the northeast, and a few Eromanga geological region-influenced catchments in the northwest of the study area. This study constitutes the largest Sr isoscape in Australia to date and

395   demonstrates the usefulness of $^{87}Sr/^{86}Sr$ analysis as a tool for understanding geological and geomorphological processes that may impact natural resources management of soil, water, and agriculture. The new spatial dataset is publicly available through the Geoscience Australia portal (https://url.place.holder.html.xxxx).

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405   Australian Government. PdC publishes with the permission from the Chief Executive Officer, Geoscience Australia.

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