<u>CCAM: China</u> Catchment <u>attributes Attributes</u> and <u>meteorology for large sample study in contiguous China Meteorology dataset</u>

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Abstract.

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The lack of a complied large-scale catchment characteristics dataset is a key obstacle limiting the development of large sample hydrology research in China. We introduce the first large-scale catchment attributes and meteorological time series dataset of contiguousin China. To develop the dataset, we'We compiled diverse data sources to generate basin oriented features describing the catchment characteristics related to hydrological processes. The proposed dataset consists of catchment characteristics, including soil, land cover, climate, topography, and geology, and 29to develop the dataset. The dataset also includes catchment scale 31-year meteorological time series (from 1990 to 2018). The meteorological variables include precipitation, temperature, evapotranspiration, wind speed, ground surface temperature, pressure, humidity and sunshine duration. We also derived a daily potential 2020 for each basin. Potential evapotranspiration time series based on a modified Penman's Penman's equation, is derived for each basin. The studied 4911 catchments are 4875 catchments within contiguous China derived from digital elevation models. We analysed and organised included in the spatial variations of catchment characteristics into a series of maps. Correlation analysis between attributes was conducted. Compared to dataset covers the entire China. We introduced several new indicators describing the catchment geography and the underlying surface compared with previously proposed datasets, we derived more eatehment characteristics. The resulting in-dataset has a total of 125 catchment attributes, providing a complete description of the catchments. Besides, we propose Normal Camels YR, a hydrological. The proposed dataset covering also includes a separate HydroMLYR dataset containing standardized weekly averaged streamflow for 102 basins of the Yellow River basin with normalized streamflow observations. Basin. The standardized streamflow data should be able to support machine learning hydrology research in the Yellow River Basin. The proposed dataset provides numerous opportunities for comparative hydrological research, such as examining the difference in hydrological behaviours across different catchments and building general rainfall runoff modelling frameworks for many eatchments instead of limited to a few. The dataset is is freely available via http://doi.org/10.5281/zenodo.4704017 for community use. We will open-source the complementat http://doi.org/10.5281/zenodo.5137288. In addition, the accompanying code for generating the dataset such that the user can generate meteorological series and catchment attributes is freely available at https://github.com/haozhen315/CCAM-China-Catchment-Attributes-and-Meteorology-dataset, supporting

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the generation of catchment characteristics for any watershed within contiguous China.custom basin boundaries. Complied data for the 4911 basins covering the entire China and the open-sourced code should be able to support the study of any arbitrary basins instead of being limited to only a few basins.

1 Introduction

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Studying a large set of catchments often provides insights that cannot be obtained when looking at a single or few catchments (Coron, Andreassian et al. 2012, Kollat, Reed et al. 2012, Newman, Clark et al. 2015, Lane, Coxon et al. 2019), The hydrologic eyele consists of many sub-processes, including evaporation from the ocean, raindrop Rainfall, interception, evaporation and evapotranspiration, groundwater flow, subsurface flow and surface runoff, infiltration, etc. are the main components of the terrestrial hydrological cycle. These processes are affected by the nature of the catchment, such as the ability of the soil to hold water. Catchment attributes such as soil characteristics, land cover characteristics and climate indices influence the water movement and storage in these sub processes of the catchment such that hydrologic behaviours can vary across catchments (van Werkhoven, Wagener et al. 2008). The same hydrological model may not be applicable in another basin. However, by examining a large sample of catchments, it is possible for the hydrological model to learn the similarities and differences of hydrological behaviours across catchments. For example, predictionStudying a large set of terrestrial catchments often provides insights that cannot be obtained when looking at a single or few (Coron, Andreassian et al. 2012, Kollat, Reed et al. 2012, Newman, Clark et al. 2015, Lane, Coxon et al. 2019). For example, a calibrated model may not be applicable in a watershed with vastly different properties. However, by examining a large sample of catchments, it is possible for a datadriven model to learn the similarities and differences of hydrological behaviours across catchments (Kratzert, Klotz et al. 2019). Prediction in ungauged basins is a challenging problem present in hydrology. The central challenge is how to extrapolate hydrologic information from gauged basins to ungauged ones. Solving, solving the problem relies on understanding the similarities and differences between different catchments. However, regionally Regionally, and temporally imbalanced observations bring a difficulty to the problem. For a hydrologic model to successfully simulate the ungauged areas, it must adapt itself to the different hydrologic behaviours present in different catchments. (Kratzert, Klotz et al. 2019) Kratzert, Klotz et al. (2019) shows encoding catchment characteristics (e.g., soil characteristics, land cover, topography) into a data-driven model can teachguide the model to behave differently responding to the meteorological time series input based on different sets of static catchment attributes.

(Silberstein 2006, Shen, Laloy et al. 2018, Nevo, Anisimov et al. 2019) pointed out that largeLarge sample hydrological datasets are the foundation and key of many hydrological studies- (Silberstein 2006, Shen, Laloy et al. 2018, Nevo, Anisimov et al. 2019). The term big hydrologic data refers to all data influencing the water cycle, such as the meteorological variables, infiltration characteristics of the study area, land use or land cover types, physical and geological features of the study areacatchment, etc. Many studies cannot be carried out without are based on large-scale hydrologic data (Coron, Andreassian

et al. 2012, Singh, van Werkhoven et al. 2014, Berghuijs, Aalbers et al. 2017, Gudmundsson, Leonard et al. 2019, Tyralis, Papacharalampous et al. 2019). For hydrological research, basin-_orientated large sample datasets are of great significance. For example, comparative hydrology (de Araújo and González Piedra 2009, Singh, Archfield et al. 2014) focus on understanding how hydrological processes interact with the ecosystem, in particular, how hydrologic behaviours change under changes in the surface and sub-surface of the earth to determine to what extent hydrological predictions can be transferred from one area to another. Large-sample catchment attributes datasetdatasets provide opportunities for research studying interrelationships among catchment attributes. (Seybold, Rothman et al. 2017)Seybold, Rothman et al. (2017) studied the correlations between river junction angle with geometric factors, downstream concavity, and aridity. (Oudin, Andréassian et al. 2008)Oudin, Andréassian et al. (2008) investigates the link between land cover and mean annual streamflow based on 1508 basins representing a large hydroclimatic variety. (Voepel, Ruddell et al. 2011)Voepel, Ruddell et al. (2011) examines how
the interaction of climate and topography influences vegetation response.

Data-driven methods can best benefit from large-scale data. Data-driven approaches have shown great potential in various fields, transforming the applications in many industries (LeCun, Bengio et al. 2015). However, data driven methods, especially the deep learning-based approaches, usually require high data volumes. Limited data will cause the over-fitting (Blumer, Ehrenfeucht et al. 1987, Abu Mostafa, Magdon Ismail et al. 2012) problem. Therefore, big hydrologic data is the fundamental support for the successful deployment of powerful data driven strategies.

Traditional hydrological models have some long standing challenges, such as the inability to capture hydrological processes' mechanism complexity (Kollat, Reed et al. 2012), which is due to the structural limitations of the conceptual models. Data-driven methods are proposed to overcome some existing obstacles. Data driven strategies open a new way for researchers to acquire knowledge transforming the research pattern from hypothesis driven to data driven. (Feng, Fang et al. 2020) proposed a flexible data integration fusing various types of observations to improve rainfall runoff modelling. The research shows that combining different resources of data benefits predictions in regions with high autocorrelation in streamflow. (Wongso, Nateghi et al. 2020) developed a model predicting the state-level, per capita water uses in the United States, taking various geographic, climatic, and socioeconomic variables as input. The research also identified key factors associated with high water usage. (Mei, Maggioni et al. 2020) proposed a statistical framework for spatial downscaling to obtain hyper-resolution precipitation data. The results show improvements compared with the original product. (Brodeur, Herman et al. 2020) applied machine learning techniques, namely bootstrap aggregation and cross validation, to reduce overfitting in reservoir control policy search. (Ni and Benson 2020) proposed an unsupervised machine learning method to differentiate flow regimes and identify capillary heterogeneity trapping, showing the promise of machine learning methods for analysing large datasets from coreflooding experiments. (Legasa and Gutiérrez 2020) propose to apply Bayesian Network for multisite precipitation occurrence generation. The proposed methodology shows improvements for existing methods.

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World-wide data sharing has become a trend (Wickel, Lehner et al. 2007, Ceola, Arheimer et al. 2015, Blume, van Meerveld et al. 2018, Wang, Chen et al. 2020), and the amounts of hydrologic data available are ever-increasing. However, these data typically came from different providers and are compiled in various formats. For example, ASTGTM (Abrams, Crippen et al. 2020) provides a global digital elevation model; GliM (Hartmann and Moosdorf 2012) includes rock types data globally; MODIS provides data products (Knyazikhin 1999, Didan 2015, Myneni, Knyazikhin et al. 2015, Running, Mu et al. 2017, Sulla-Menashe and Friedl 2018) describing features of the land and the atmosphere derived from remote sensing observations; (Yamazaki, Ikeshima et al. 2019) yamazaki, Ikeshima et al. (2019) provides a global flow direction map at three arc-second resolution; HydroBASINS (Lehner 2014) provides basin boundaries at different scales globally; and GDBD (Masutomi, Inui et al. 2009) provides basin boundaries with geographic attributes; GLHYMPS (Gleeson, Moosdorf et al. 2014) provides a global map of subsurface permeability and porosity; SoilGrids250m (Hengl, Mendes de Jesus et al. 2017) dataset provides global numeric soil properties. Local government agencies often hold meteorological data such as precipitation and evaporation, and the amount of this data is also growing, however, data transparency has still been a problem (Viglione, Borga et al. 2010). The

However, the data mentioned above are rarely spatially aggregated to the catchment-scale, making it difficult for researchers to use these data. Properly pre-processed and formatted datasets on a large scale are of great importance for the hydrology research. Searching for appropriate data sources, pre-processing, and formatting often consumes a lot of researchers' time. In some cases, individual research groups either do not know where to obtain the appropriate data or cannot properly process the data to receive the desired format. In summary, although data sharing is being advocated in the community, it is usually difficult for the public to obtain the required data, either because there are not enough observations or because of the difficulties in the data processing.

In summary, both data driven and traditional hydrological research need diverse hydrologic datasets to learn the generalisation capability from one area to another. For a model to adapt to various behaviours in different catchments, the dataset must be large enough to represent the complex heterogeneity presented in the natural hydrologic system. Although data sharing is being advocated in the community, it is usually difficult for the public to obtain certain data such as meteorological data and streamflow observations, either because there are not enough observations or because there are no open access permissions.

Recently, there are efforts (Addor, Newman et al. 2017, Alvarez-Garreton, Mendoza et al. 2018, Chagas, Chaffe et al. 2020, Coxon, Addor et al. 2020) compiling different types of data sources to form large scale hydrological datasets. These four collected datasets cover the continental United States, Chile, Brazil, and Great Britain. (Addor, Do et al. 2020) reviewed these datasets and discussed the guidelines for producing large sample hydrological datasets and the limitations of the currently proposed datasets. The CAMELS dataset has been used to support a lot of research. Based on CAMELS, (Kratzert, Klotz et

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¹⁻https://asterweb.jpl.nasa.gov/gdem.asp

al. 2018) built a Long Short-Term Memory (LSTM) network for rainfall-runoff modelling, showing that one model can predict the discharge for a variety of catchments. (Knoben, Freer et al. 2019) compared metrics used in hydrology based on simulations on many basins. (Tyralis, Papacharalampous et al. 2019) studied the relationship between the shape parameter and basin attributes based on the sizeable basin-oriented dataset.

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However, there is no large-scale compilation of hydrological datasets in contiguous China. An alternative is on a global scale, the HydroATLAS (Linke, Lehner et al. 2019) dataset. However, since it is on a world-wide scale, compared with other datasets constructed for regions, the dataset lacks many attributes and is not built according to the CAMELS standards. Besides, the elimatic data is not up to date (1950-2000), and the derivation of climatic data lacks ground surface observations inputs, such that the data quality is not guaranteed.

Therefore, researchers still need to do repetitive works to compile data from different sources such as obtaining historical meteorological data (temperature, rainfall, evapotranspiration) of a catchment in contiguous China. Inspired by (Addor, Newman et al. 2017), in this paper, we present a catchment scale hydrologic dataset compiling a wide variety of hydrological data, including basin topography, climate indices, land cover characteristics, soil characteristics and geological characteristics covering contiguous China.

to compile different types of data sources forming large scale hydrological datasets. These four collected datasets cover the continental United States, Chile, Brazil, and Great Britain. Addor, Do et al. (2020) reviewed these datasets and discussed the guidelines for producing large-sample hydrological datasets and the limitations of the currently proposed datasets. The static properties of 671 river basins in the United States are calculated by CAMELS (Addor, Newman et al. 2017), which is an extension of a previously proposed hydrometeorological data set (Newman, Clark et al. 2015). Unfortunately, it is impossible to publish streamflow data in China for the time being. The CAMELS dataset has been used to support a lot of research. For example, Knoben, Freer et al. (2019) compared metrics used in hydrology based on simulations on many basins. Tyralis, Papacharalampous et al. (2019) studied the relationship between the shape parameter and basin attributes based on the sizeable basin-oriented dataset.

There is currently no compilation of China-specific catchment attributes datasets. An alternative, the HydroATLAS (Linke, Lehner et al. 2019) dataset, which is on a global scale, is basically performing zonal statistics on the source data. HydroATLAS lacks many indicators which need derivations based on the source data, such as rainfall seasonality, the fraction of precipitation falling as snow, basin shape factors and root depth distributions. What's worse, the meteorological data is only up to 2000, which is outdated.

In summary, a lack of a complied catchment attributes dataset is a key obstacle limiting the development of large sample hydrology research in China. Inspired by (Addor, Newman et al. 2017), we complied multiple data sources, including basin topography, climate indices, land cover characteristics, soil characteristics and geological characteristics. Different from

(Addor, Newman et al. 2017), the catchments included in the dataset covers the entire study area, instead of being limited to a few.

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The proposed dataset is the first dataset providing catchments meteorological time series and catchments attributes of contiguous China. We compiled and named the dataset following most standards of the previously proposed datasets. Unlike CAMELS and CAMELS CL, catchments in the proposed dataset are not selective. Instead, the The dataset consists of all generated basins derived basin boundaries from the Digital Elevation Model (DEM), based on) which came from the Global Drainage Basin Dataset (Masutomi, Inui et al. 2009). The Global Drainage Basin Dataset (GDBD) is derived at high-resolution (100m-1km) and has a good geographic agreement with existing global drainage basin data in China². Besides, an essential feature of the. In addition, previously proposed dataset is that it provides a complete description of the catchment, rather than an abstraction. For example, both CAMELS and CAMELS CL only datasets (Addor, Newman et al. 2017, Alvarez-Garreton, Mendoza et al. 2018, Chagas, Chaffe et al. 2020, Coxon, Addor et al. 2020) report only the most frequent and second most frequent-catchment land cover and lithology types. Instead, the proposed dataset CCAM calculates the proportion proportions of eachall land cover and lithology type for each catchment to serve data driven research better. We also introduced many more climate characteristics and soil characteristics to support more diverse potential research types.

Researchers from different places can use the proposed dataset in conjunction with their streamflow data, simplifying organising and compiling various data resources, which is usually repetitive work. The proposed dataset is undoubtedly the most comprehensive catchment attributes and meteorological time series dataset in contiguous China and is suitable for multipurpose data driven research. The dataset consists of basin boundaries in the shapefile format, computed catchment attributes of climate, land cover, soil, topography and lithology and 29 year meteorological time series. Table 1 compares the number of static attributes between CAMELS, CAMELS BR, and the proposed dataset.

In addition to the basin-wise attributes provided in CCAM, we propose HydroMLYR, a hydrology dataset for machine learning research in the Yellow River Basin providing weekly averaged standardized streamflow data for 102 basins in the Yellow River Basin (YRB). HydroMLYR is proposed to support machine learning hydrology research at YRB. Traditional hydrological models have some long standing challenges, such as the inability to capture hydrological processes' mechanism complexity (Kollat, Reed et al. 2012), which is due to the structural limitations of the conceptual models. Data-driven strategies represented by machine learning are proposed to overcome some existing obstacles and they open a new way for researchers to acquire knowledge transforming the research pattern from hypothesis-driven to data-driven. Feng, Fang et al. (2020)

² In this study, gauge streamflow measurements are not available in areas other than the Yellow River such that it is infeasible to specify a gauge location for generating the basin boundary for most of the areas. Streamflow measurements have strict redistribution policy; however, local research institutions have their streamflow measurements for hydrological research, the proposed dataset can used in conjunction with the streamflow data of researchers in various places.

proposed a flexible data integration fusing various types of observations to improve rainfall-runoff modelling. The research shows that combining different resources of data benefits predictions in regions with high autocorrelation in streamflow. Wongso, Nateghi et al. (2020) developed a model predicting the state-level, per capita water uses in the United States, taking various geographic, climatic, and socioeconomic variables as input. The research also identified key factors associated with high water usage. Mei, Maggioni et al. (2020) proposed a statistical framework for spatial downscaling to obtain hyper-resolution precipitation data. The results show improvements compared with the original product. Brodeur, Herman et al. (2020) applied machine learning techniques, namely bootstrap aggregation and cross-validation, to reduce overfitting in reservoir control policy search. Ni and Benson (2020) proposed an unsupervised machine learning method to differentiate flow regimes and identify capillary heterogeneity trapping, showing the promise of machine learning methods for analysing large datasets from coreflooding experiments. Legasa and Gutiérrez (2020) propose to apply Bayesian Network for multisite precipitation occurrence generation, and the proposed methodology shows improvements for existing methods. The proposed data set can be used to develop or verify machine learning models in the YRB.

The paper is organized as follows: Section 2 describes the study area. Section 3-7 describes the five classes of the computed catchment attributes. In section 3-7, each unit follows the same structure: first introduce the meaning and significance of each added feature and data source used, then describe the variables' spatial variability if necessary. Section 8 describes the proposed catchment_scale meteorological forcing time series. Section 9 introduce the Normal Camels YRHydroMLYR dataset, which provides normalized streamflow measurements for 102 catchments of Yellow River. Section 10 describes the code and data availability. Section 11 presents the concluding remark.

In summary, our contributions are as follows:

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- (1) The proposed dataset is the first large scale dataset containing catchment scale meteorological time series of contiguous China, which is the basis for many hydrological studies.
- (2) We present the first basin oriented static attributes dataset in contiguous China.
- (3) We introduce several new catchment characteristics providing a complete description of the catchment compared with the previously proposed datasets such that the proposed dataset is prepared for potential hydrological studies.
- (4) We offer a self-contained dataset covering 102 basins of the Yellow River basin with normalized runoff observation supporting many potential studies.
- (5) We will open source the code for generating the dataset such that the user can generate a dataset for any watershed within contiguous China.

Table 1 Number of computed attributes in CAMELS, CAMELS BR and the proposed dataset.

| Attribute class | CAMELS(A | 17) | CAMELS-BR | Ours |
|-------------------------|----------|----------------|---------------|------|
| Location and | 9 | # | 12 | |
| topography | | | | |

| Geology | 7 | 7 | 18 |
|-----------------------|---------------------|---------------|----------------|
| Soil | ## | 6 | 54 |
| Land cover | <u> </u> | # | 22 |
| Climatic indices | # | 13 | 17 |
| Human intervention | not | 4 | ⊋ |
| indices | computed | | |
| Total | 46 | 52 | 125 |
| | | | |

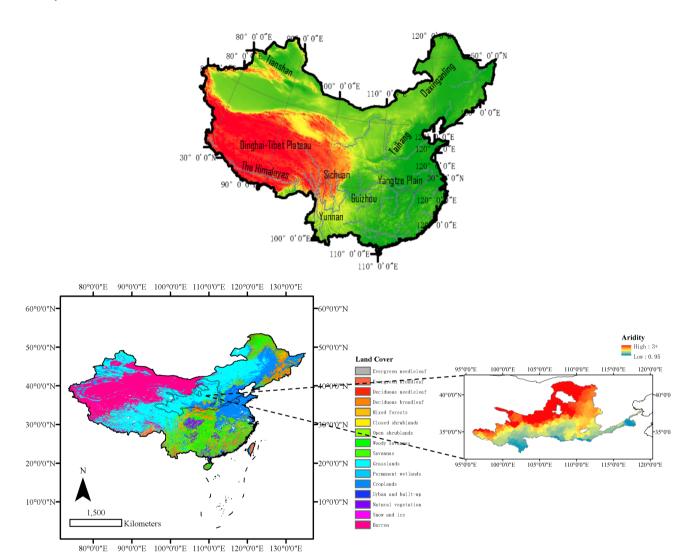
225 Table 2 Summary of basin daily discharge and foreing data in CAMELS, CAMELS-BR and the proposed dataset.

| G L I CET C | | GALLET G DD | |
|----------------------|---|---|---|
| CAMELS | | CAMELS-BR | Ours |
| available | available | available | |
| available | available | available | |
| available | not | available | |
| | available | | |
| available | not | not available | |
| | available | | |
| net | not | available | |
| available | available | | |
| available | not | available | |
| | available | | |
| available | not | not available | |
| | available | | |
| not | not | available | |
| available | available | | |
| available | not | available | |
| | available | | |
| not | available | available | |
| available | | | |
| not | available | available | |
| available | | | |
| available | | available | partially available (see Section 9) |
| | available | available | available |

2 Study area

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(a)

Figure 1. Overview of the study area. The study area covers a wide range of latitude and longitude, from 18.2° N to 52.3° N, and from 76.0° E to 134.3° E. (a) The main geographical features map of contiguous China. China is mountainous; mountains and hills occupy two-thirds of the area. (b) The distribution map of the delimited eatehments based on the ASTER DEM, the eatehments studied are all catchment areas delimited from the DEM, covering contiguous China, with 4875 catchments, most of which are 2000 to 5000 square kilometres.

: Left: Study area of CCAM and the distribution of land cover types. The studied basins cover the whole of China. Right: Study area of HydroMLYR and the distribution of aridity (PET/P) index. YRB is a generally arid area. The data set provided can be used as a good sample for studying hydrology in arid regions.

The study area corresponds to contiguous the whole of China; (Fig. 1), with diverse climate and terrain characteristics, spanning from 18.2° N to 52.3° N and 76.0° E to 134.3° E. Mountains, plateaus, and hills account for about two-thirds of areas of contiguous China, and the remaining are basins and plains. China's topography is like a three-level ladder, high in the west and low in the east. The Qinghai-Tibet Plateau, the highest plateau globally, located in the west of contiguous China, with a mean elevation of over 4000 meters, is the first step of China's topography. The Xinjiang region, the Loess Plateau, the Sichuan Basin, and the Yunnan-Guizhou Plateau to the north and east are the second step of China's topography. The mean sea level here is between 1000 to 2000 meters. Plains and hills dominate the east of the Daxinganling-Taihang Mountain to the coastline, the third step of contiguous China. The elevation of this step descends to 500-1,000 meters. To better characterize the studied catchments, we have derived various attributes. Table 1 compares the number of derived attributes between several proposed datasets.

Table 1: Number of computed attributes in CAMELS, CAMELS-BR and CCAM.

| Attribute class | CAMELS(A17) | CAMELS-BR | <u>CCAM</u> |
|----------------------------|--------------|-----------|-------------|
| Location and topography | <u>9</u> | <u>11</u> | <u>12</u> |
| Geology | <u>7</u> | <u>7</u> | <u>18</u> |
| Soil | <u>11</u> | <u>6</u> | <u>54</u> |
| <u>Land cover</u> | <u>8</u> | <u>11</u> | <u>22</u> |
| Climatic indices | <u>11</u> | <u>13</u> | <u>17</u> |
| Human intervention indices | not computed | <u>4</u> | <u>2</u> |
| <u>Total</u> | <u>46</u> | <u>52</u> | <u>125</u> |

In—contiguous China, precipitation and temperature vary significantly in different places, forming a diverse climate environment. According to the Köppen Climate Classification System, from northwest to southeast, China's climate gradually evolves from Cold desert (BW_k) climate, Tundra (ET) climate, Warm and temperate continental (D_{fa} and D_{wb}) climate to Humid subtropical (C_{wa}) climate and Warm oceanic (C_{fa}) climate. From the perspective of temperature zones, there are tropical, subtropical, warm temperate, medium temperate, cold temperate and Qinghai-Tibet Plateau regions, and there are humid regions, semi-humid regions, semiarid regions, and arid regions from the perspective of wet and dry zones. Moreover, the same temperature zone can contain different dry and wet zones. Therefore, there will be differences in heat and wetness in the same climate type. The complexity of the terrain makes the climate even more complex and diverse. Besides, China has a wide range of regions affected by the alternating winter and summer monsoons. Compared with other parts of the world at the same latitude, these areas have low winter temperatures, high summer temperatures, significant annual temperature differences, and concentrated precipitation in summer. The cold and dry winter monsoon occurs in Asia's interior, far away from the ocean. Under its influence, winter rainfall in most parts of China is low, accompanied by low temperature. The summer monsoon is

warm and humid, coming from the Pacific Ocean and the Indian Ocean. Under its influence, precipitation generally increases. Table 2 compares the provided forcing variables in CAMELS, CAMELS-BR and CCAM.

Table 2: Summary of forcing variables provided in CAMELS, CAMELS-BR and CCAM.

| Forcing data class | CAMELS | CAMELS-BR | CCAM |
|------------------------------|----------------------|----------------------|----------------------|
| <u>Temperature</u> | <u>available</u> | available | available |
| <u>Precipitation</u> | <u>available</u> | <u>available</u> | <u>available</u> |
| Solar radiation | <u>available</u> | <u>not available</u> | <u>available</u> |
| Day length | <u>available</u> | <u>not available</u> | <u>not available</u> |
| Sunshine hours | not available | not available | <u>available</u> |
| <u>Humidity</u> | <u>available</u> | not available | <u>available</u> |
| Snow water equivalent | <u>available</u> | <u>not available</u> | <u>not available</u> |
| Wind velocity | <u>not available</u> | <u>not available</u> | <u>available</u> |
| Ground surface pressure | <u>available</u> | <u>not available</u> | <u>available</u> |
| Observed evaporation | <u>not available</u> | <u>available</u> | <u>available</u> |
| Potential evapotranspiration | not available | <u>available</u> | <u>available</u> |

Table 3: Summary table of catchment attributes available in the proposed dataset.

| Attribute class | Attribute name | Description | Unit | Data source |
|-----------------|----------------|---------------------------------|-------|-------------------------------------|
| Climate indices | pet_mean | mean daily pet (Penman-Monteith | mm d | (Subramanya 2013) |
| (computed for 1 | | equation) | 1 | |
| Oct 1990 to 30 | evp_mean | mean daily evaporation | mm d- | SURF_CLI_CHN_MUL_DAY3F ³ |
| Sep 2018) | | (observations) | 1 | |
| | gst_mean | mean daily ground surface | °C | • |
| | | temperature | | |
| | pre_mean | mean daily precipitation | mm d | - |
| | | | 1 | |
| | prs_mean | mean daily ground surface | hPa | • |
| | | pressure | | |
| | rhu_mean | mean daily relative humidity | - | • |
| | ssd_mean | mean daily sunshine duration | h | • |

³ http://data.cma.cn/data/cdcdetail/dataCode/SURF CLI CHN MUL DAY.html

| tem_mean | mean daily temperature | °C |
|------------------|--|--------------------|
| win_mean | mean daily wind speed | m s ⁻¹ |
| p_seasonality | seasonality and timing of | - |
| | precipitation (estimated using sine | |
| | curves to represent the annual | |
| | temperature and precipitation | |
| | cycles, positive [negative] values | |
| | indicate that precipitation peaks in | |
| | summer [winter], values close to 0 | |
| | indicate uniform precipitation | |
| | throughout the year) | |
| high_prec_freq | frequency of high-precipitation | d yr ⁻¹ |
| | days (≥ 5 times mean daily | |
| | precipitation) | |
| high_prec_dur | average duration of high- | d |
| | precipitation events (number of | |
| | consecutive days \geq 5 times mean | |
| | daily precipitation) | |
| high_prec_timing | season during which most high- | season |
| | precipitation days (≥ 5 times | |
| | mean daily precipitation) occur | |
| low_prec_freq | frequency of dry days (< 1mm d ⁻¹) | d yr ⁻¹ |
| low_prec_dur | average duration of dry periods | d |
| | (number of consecutive days < 1 | |
| | mm d ⁻¹) | |
| low_prec_timing | season during which most dry days | season |
| | (< 1 mm d ⁻¹) occur | |
| frac_snow_daily | fraction of precipitation falling as | - |
| | snow (for days colder than 0 °C) | |
| p_seasonality | seasonality and timing of | - |
| | precipitation, positive [negative] | |
| | values indicate that precipitation | |
| | peaks in summer [winter], values | |
| | | |

| | | close to 0 indicate uniform | |
|-----------------|-------------------|---|---------------------------------|
| | | precipitation throughout the year | |
| Geological | geol_porosity | subsurface porosity - | (Gleeson, Moosdorf et al. 2014) |
| characteristics | geol_permeability | subsurface permeability (log-10) m ² | |
| | ig | fraction of the catchment area - | (Hartmann and Moosdorf 2012) |
| | | associated with ice and glaciers | |
| | pa | fraction of the catchment area - | |
| | | associated with acid plutonic rocks | |
| | sc | fraction of the catchment area - | <u> </u> |
| | | associated with carbonate | |
| | | sedimentary rocks | |
| | su | fraction of the catchment area - | <u> </u> |
| | | associated with unconsolidated | |
| | | sediments | |
| | sm | fraction of the catchment area - | |
| | | associated with mixed | |
| | | sedimentary rocks | |
| | vi | fraction of the catchment area - | <u> </u> |
| | | associated with intermediate | |
| | | volcanic rocks | |
| | mt | fraction of the catchment area - | <u> </u> |
| | | associated with metamorphic | |
| | SS | fraction of the catchment area - | <u> </u> |
| | | associated with siliciclastic | |
| | | sedimentary rocks | |
| | pi | fraction of the catchment area - | |
| | | associated with intermediate | |
| | | plutonic rocks | |
| | va | fraction of the catchment area - | <u> </u> |
| | | associated with acid volcanic | |
| | | rocks | |
| | wb | fraction of the catchment area - | <u> </u> |
| | | associated with water bodies | |

| | pb | fraction of the catchment area - associated with basic plutonic rocks | - | |
|----------------------------|------------------------------|---|--------------|----------------------------------|
| | vb | fraction of the catchment area - associated with basic volcanic rocks | - | |
| | nd | fraction of the catchment area - associated with no data | - | |
| | ру | fraction of the catchment area - associated with pyroclastic | - | |
| | ev | fraction of the catchment area - associated with evaporites | | |
| Land cover characteristics | lai_max | maximum monthly mean of the leaf area index (based on 12 monthly means) | - | (Myneni, Knyazikhin et al. 2015) |
| | lai_diff | and minimum monthly mean of the leaf area index (based on 12 monthly means) | - | |
| | ndvi_mean | mean normalized difference - vegetation index (NDVI) | - | (Didan 2015) |
| | root_depth_50 | root depth (percentiles=50% rextracted from a root depth distribution based on IGBP land cover) | m | Eq. 2 and Table 2 in (Zeng 2001) |
| | root_depth_99 | root depth (percentiles=99% restracted from a root depth distribution based on IGBP land cover) | m | |
| | evergreen needleleaf tree | catchment area fraction covered - by evergreen needleleaf tree | | (Sulla-Menashe and Friedl 2018) |
| | evergreen broadleaf tree | catchment area fraction covered - by evergreen broadleaf tree | - | |

| | deciduous | catchment area fraction covered | - | |
|-----|--------------------|---------------------------------|--------|------------------------------|
| | needleleaf tree | by deciduous needleleaf forests | | |
| | deciduous | catchment area fraction covered | - | • |
| | broadleaf tree | by deciduous broadleaf tree | | |
| | mixed forest | catchment area fraction covered | - | • |
| | | by mixed forest | | |
| | closed shrubland | catchment area fraction covered | - | • |
| | | by closed shrubland | | |
| | open shrubland | catchment area fraction covered | - | • |
| | | by open shrubland | | |
| | woody savanna | catchment area fraction covered | - | • |
| | | by woody savanna | | |
| | savanna | catchment area fraction covered | - | • |
| | | by savanna | | |
| | grassland | catchment area fraction covered | - | <u>.</u> |
| | | by grassland | | |
| | permanent | catchment area fraction covered | - | • |
| | wetland | by permanent wetland | | |
| | cropland | catchment area fraction covered | - | • |
| | | by cropland | | |
| | urban and built-up | catchment area fraction covered | - | • |
| | land | by urban and built-up land | | |
| | cropland/natural | catchment area fraction covered | - | • |
| | vegetation | by cropland/natural vegetation | | |
| | snow and ice | catchment area fraction covered | - | • |
| | | by snow and ice | | |
| | barren | catchment area fraction covered | - | • |
| | | by barren | | |
| | water bodies | catchment area fraction covered | - | • |
| | | by water bodies | | |
| 7, | basin_id | drainage basin identifiers | - | (Masutomi, Inui et al. 2009) |
| and | pop | population | people | |
| | | | | |

Topography, location,

| Human | | | people | |
|--------------|-------------------|---|-------------------|------------------------------|
| intervention | pop_dnsty | population density | km ⁻² | |
| | lat | mean latitude | °N | - |
| | lon | mean longitude | °E | - |
| | elev | mean elevation | M | - |
| | area | catchment area | km ² | - |
| | | | m km ⁻ | (Horn 1981) |
| | slope | mean slope | 1 | |
| | length | The length of the mainstream | kmKm | (Subramanya 2013) |
| | | measured from the basin outlet to | | |
| | | the remotest point on the basin | | |
| | | boundary. The mainstream is | | |
| | | identified by starting from the | | |
| | | basin outlet and moving up the | | |
| | | catchment. | | |
| | form factor | catchment area / (catchment | - | <u>-</u> |
| | | length) ² | | |
| | shape factor | (catchment length) ² / catchment | - | - |
| | | area | | |
| | compactness | perimeter of the catchment / | - | - |
| | coefficient | perimeter of the circle whose area | | |
| | | is that of the basin | | |
| | circulatory ratio | catchment area / area of circle of | - | - |
| | | catchment perimeter | | |
| | elongation ratio | diameter of circle whose area is | - | - |
| | | basin area / catchment length | | |
| Soil | pdep | soil profile depth | cm | (Shangguan, Dai et al. 2013) |
| | clay | percentage of clay content of the | % | - |
| | | soil material | | |
| | sand | percentage of sand content of the | % | - |
| | | soil material | | |
| | por | porosity | cm ³ | - |
| | | | cm ⁻³ | |

| silt | percentage of silt content of the | % | |
|------------------------|------------------------------------|--------------------|--------------------------------|
| | soil material | | |
| grav | rock fragment content | % | • |
| som | soil organic carbon content | % | • |
| log_k_s4F ⁴ | log-10 transformation of saturated | cm d ⁻¹ | (Dai, Xin et al. 2019) |
| | hydraulic conductivity | | |
| theta_s ⁴ | saturated water content | cm ³ | • |
| | | cm ⁻³ | |
| tksatu ⁴ | thermal conductivity of unfrozen | W m ⁻¹ | • |
| | saturated soils | K^{-1} | |
| bldfie ⁴ | bulk density | kg m ⁻³ | (Hengl, Mendes de Jesus et al. |
| cecsol ⁴ | cation-exchange capacity | cmol+ | 2017) |
| | | kg ⁻¹ | |
| orcdrc ⁴ | organic carbon content | g kg ⁻¹ | • |
| phihox ⁴ | pH in H2O | 10-1 | • |
| bdticm | depth to bedrock | cm | • |
| | | | |

3 Climate Climatic indices

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Meteorological raw Raw meteorological data wasis provided by the China Meteorological Data Network, released as the SURF_CLI_CHN_MUL_DAY (V3.0) dataset⁵, which provides complete variable types and the longest period (1951-20182020) of meteorological time series ofin China. The SURF_CLI_CHN_MUL_DAY product includes site observations of pressure, temperature, relative humidity, precipitation, evaporation, wind speed, sunshine duration, and ground surface temperature. The summary is presented in Table 4. (Table 4). The Inverse distance weighting method is used for interpolating the site observations. Climate indices are then obtained by taking the average of the catchment-scale extraction from the interpolated raster. To ensure data quality, we choseuse the latter 2931-year record (from 1990 to 20182020) to construct the dataset since sites' distribution was sparse in the early days (Fig. 2). We computed more climatic characteristics compared with other datasets (Table 2). These characteristics have critical potential effects on theyariables are useful in hydrological processes modelling; for example, wind speed can affect actual evapotranspiration. To be consistent with the CAMELS (Addor, Newman et al. 2017), we also determined all climatic attributes (Woods 2009) provided in the CAMELS dataset. The As a

⁴ The data source contains multi-layer soil data, soil characteristics for all layers are determined.

⁵ SURF CLI CHN MUL DAY is freely available for global researchers.

result, the proposed dataset provides more meteorological variables and longer time series (1990-20182020) than CAMELS and CAMELS-CL. A summary of the computed Climate derived climate indices is presented in Table 3. The national distribution distributions of meteorological attributes of catchments is the climate indicators are shown in Fig. 3. Fig. 3.

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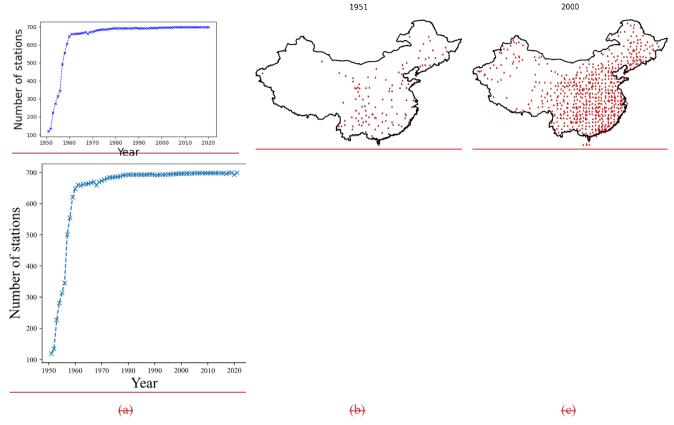
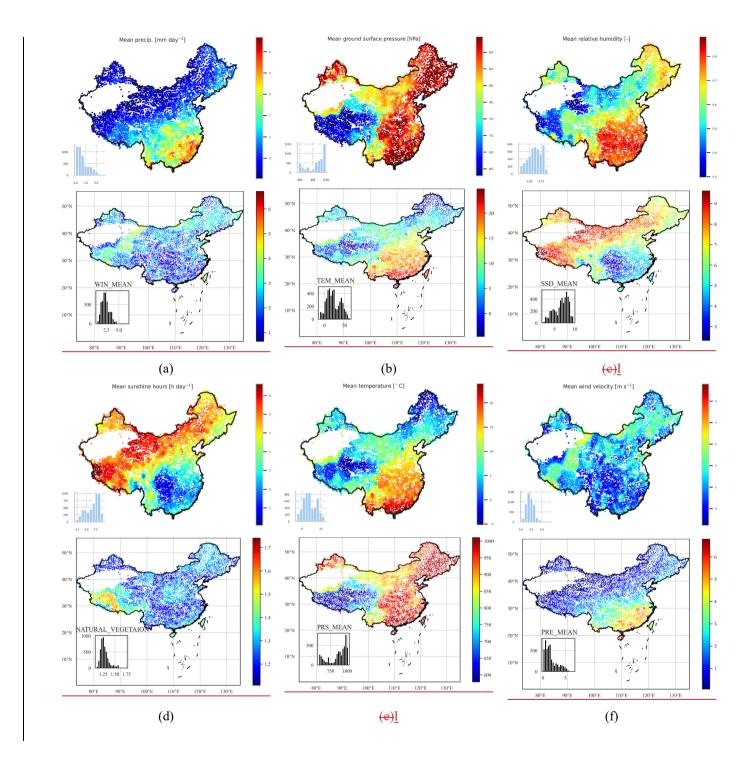
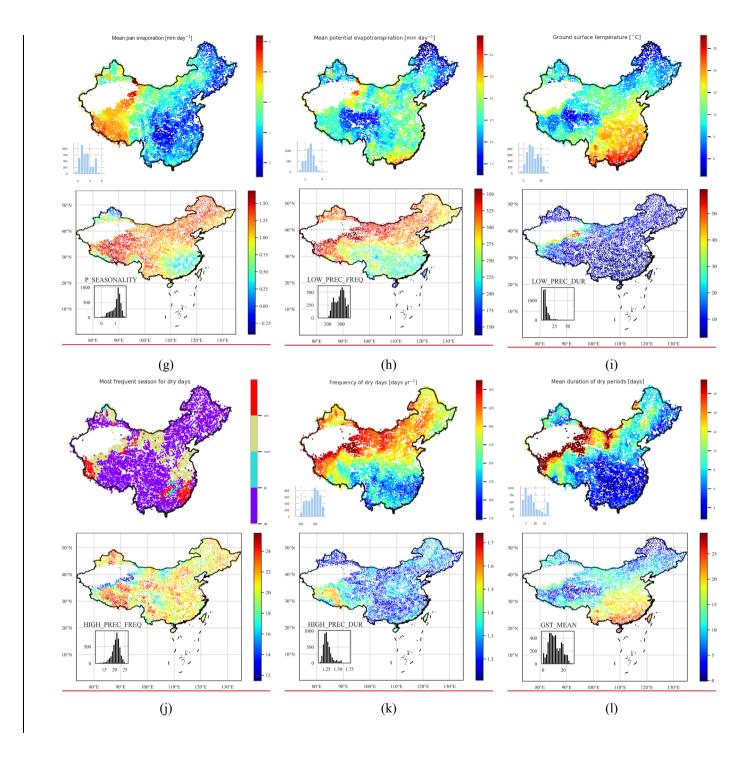


Figure 2. Overview of changes: Changes in the number and distribution of meteorological stations in China. (a) The number of meteorological stations varies with the year. There were only 119 stations in 1951. This number increased rapidly from 1951 to the early 1960s, and the number of stations remained stable after 2000. (b) Distribution map of China's meteorological stations in 1951. (c) Distribution map of China's meteorological stations in 2000 To ensure the data quality, we used the latter 31-year records (from 1990 to 2020) to construct the dataset.





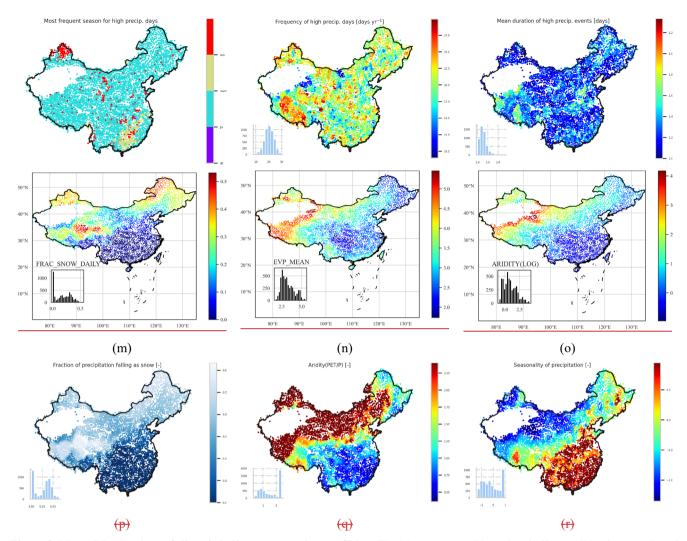


Figure 3. Maps: Distributions of climatic indices over contiguous. China. The histograms and bar plots indicate All basins are plotted in the numbersame size. When extreme values of catchments (out of 4875) in each bin or category a variable affect visualization (cause most areas to have the same colour), the log values are used for visualization.

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The instruments for measuring <u>potential</u> evaporation were updated from 2000 to 2005. Early observations can be multiplied by a correction coefficient to approximate the new tools. However, the coefficient varies across stations making the approach infeasible. To complement this, we calculated potential evapotranspiration (PET) based on a modified Penman's Equation (see Appendix A) and other observed meteorological variables, providing a series of consistent <u>evapotranspiration</u> <u>estimation potential evaporation estimations for reference</u>.

The average daily precipitation in contiguous China is highest in the southeast and lowest in the northwest. It is also higher in the coastal areas than in the interior land. Ground surface pressure is positively correlated with elevation, the highest in the Qinghai-Tibet Plateau and the lowest in the Southeast Plain. The average relative humidity is generally positively correlated

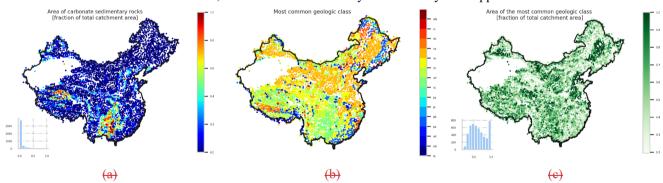
with precipitation; they are also higher in some forested areas, such as the Taihang Mountains and Daxingan Mountains. The Qinghai-Tibet Plateau has the lowest average temperature, and the southern coastal area has the highest. A distinctive feature of the distribution of wind speed is the high wind speed in mountainous areas. The highest wind speed occurs in the southeast coastal area (> 6 meters per second). Refer to Section 8 for a detailed description of the proposed catchment-scale meteorological time series dataset of contiguous China.

4 Geology

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To describe the lithological characteristics of each catchment, we used the same two global datasets as CAMELS, Global Lithological Map (GLiMGliM) (Hartmann and Moosdorf 2012) and GLobal Hydrogeology Global Hydrogeology MaPS (GLHYMPS) (Gleeson, Moosdorf et al. 2014). Figure 4 presents the results distributions of the geological types.

310 GLiM provides a high resolution global lithological map assembled from existing regional geological maps; it has been widely used for constructing datasets (e.g. SoilGrids250m (Hengl, Mendes de Jesus et al. 2017)). However, the data quality of GLiM can vary in different spatial locations depending on the quality of the original regional geological maps. GLiM consists of three levels, the first level contains 16 lithological classes, and the additional two levels describe more specific lithological characteristics. The GLiM is represented by 1,235,400 polygons; the polygons are converted to raster format for the basin-scale lithological type statistics. For contiguous For China, the compiled regional data sources (China 1991, Xinjiang 1992, Survey 2001) have slightly lower resolutions than the GLiM target resolution (1:1 000 000). However, for a basin-scale study with a mean basin area of over 2000 km², the classification accuracy should satisfy most applications.



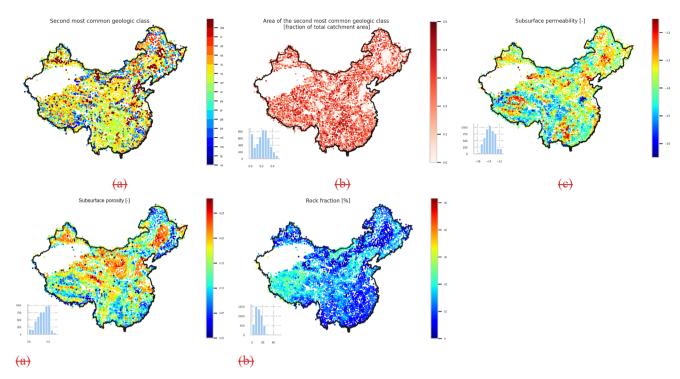
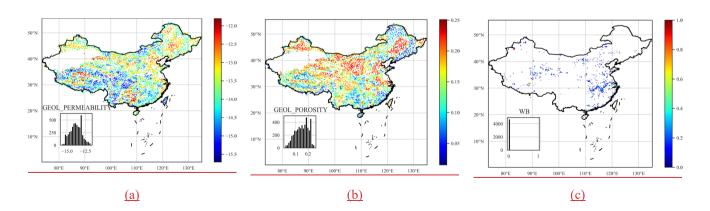
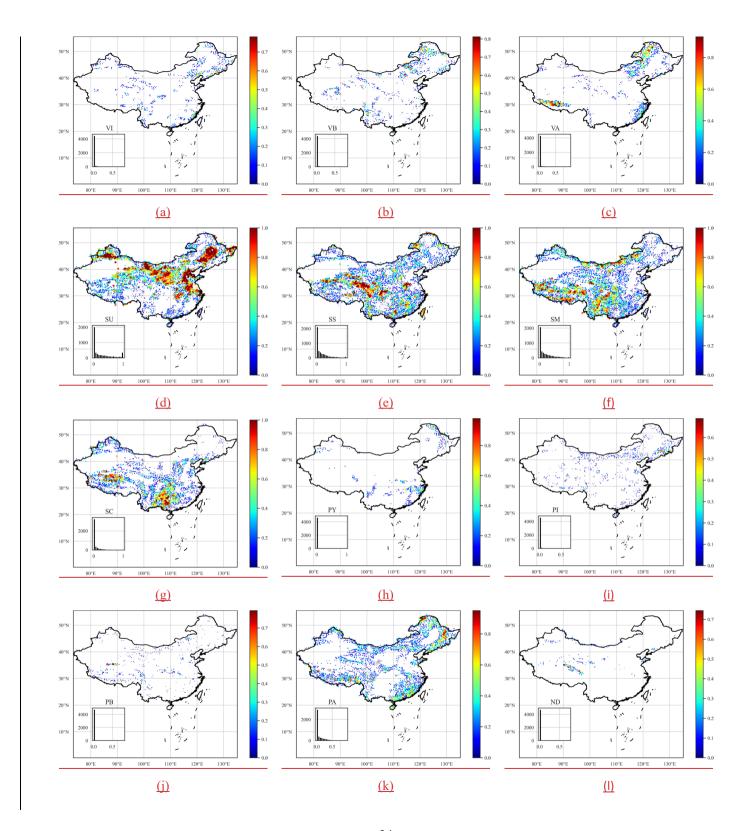


Figure 4. Maps of geological characteristics over contiguous China. The histograms indicate the number of catchments (out of 4875) in each bin.

320 Compared to CAMELS and CAMELS CL, one design consideration of the proposed dataset is that it should be more prepared for the data driven research, such that we aim to generate as many types of catchment scale data as possible since advanced data driven methods can learn the representation of inputs automatically. To this end, we determined and recorded <u>Different from CAMELS and CAMELS-CL</u>, we determined each lithological class's contribution to the catchment instead of recoding just the first and second most frequent classes. The GLiM is represented by 1,235,400 polygons; the polygons are converted to rester format for the basin scale lithological type statistics.





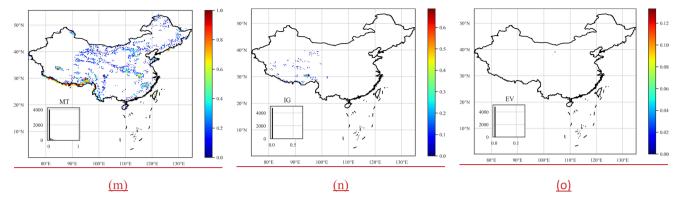


Figure 4: Distributions of geological characteristics over China. For lithologies, the plot size is scaled by the lithology proportion.

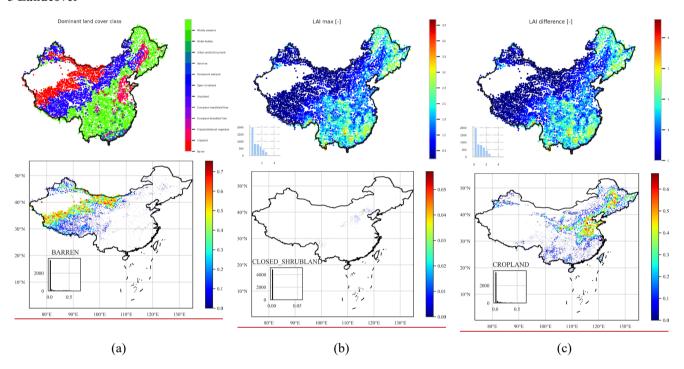
GLobal HYdrogeology MaPS (GLHYMPS) provides a global estimation of subsurface permeability and porosity, two critical characteristics for the soils' hydrological classification. Porosity and permeability influence an area's infiltration capacity. Soil with high porosity is likely to contain s amounts of water, and high permeable soil transmits water relatively quickly. Based on the high-resolution map of GLiM, which can differentiate fine and coarse-grained sediments and sedimentary rocks, GLHYMPS determined subsurface permeability depending on the different permeabilities of rock types. For the proposed dataset, we calculated the catchment arithmetic mean for porosity. Followed (Gleeson, Smith et al. 2011), the logarithmic scale geometric mean is used for representing subsurface permeability. The summary of geological characteristics is present in Table 3.

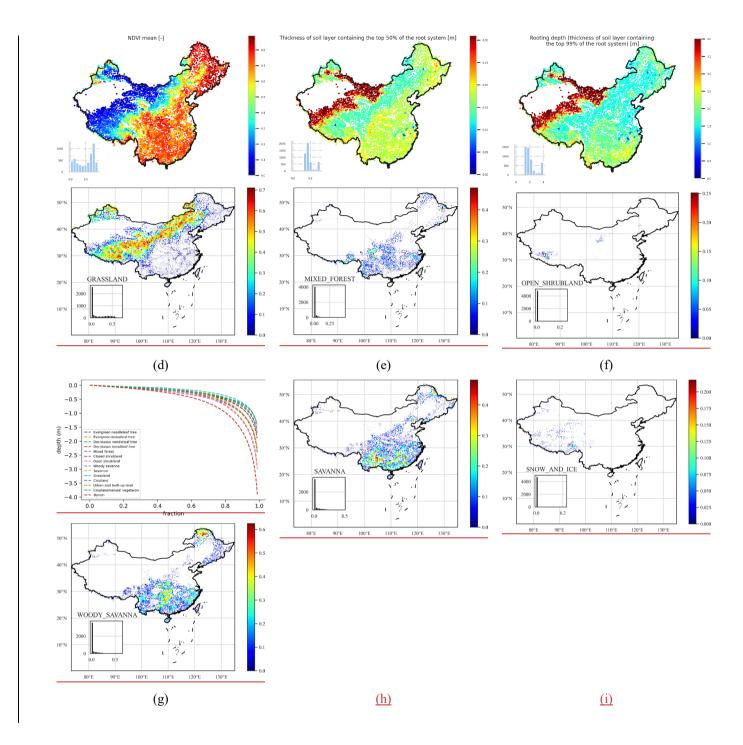
Porosity and permeability have similar distributions as geological classes. These two characteristics are highly dependent on rock properties, unconsolidated sediments, mixed sedimentary rocks, siliciclastic sedimentary rocks, carbonate sedimentary rocks, and acid plutonic rocks are the five most common geological classes in contiguous. China. Unconsolidated sediment is the most common rock type in contiguous. China, dominating 31.9% of catchments; it extends from Xinjiang to the inland of the northeast and the coastal area surrounding the Bohai Sea, due to the high proportion of unconsolidated sediments present in the rock, these areas typically have high permeability and medium porosity. Mixed sedimentary rocks are the second most common rock type in contiguous. China, accounting for 20.3% of catchments, it dominated the southern Qinghai-Tibet Plateau, western Yunnan-Guizhou Plateau, and northern Inner Mongolia. These areas typically have high porosity and low permeability. Siliciclastic sedimentary rocks dominate 17.7% of basins, mainly distributed in the northern part of the Qinghai-Tibet Plateau and the junction of the Qinghai-Tibet Plateau and the Yunnan-Guizhou Plateau; there are also some distributions in the eastern inland. These areas have low subsurface permeability and high subsurface porosity. Amongst all catchments, 9.8% of catchments are dominated by carbonate sedimentary rocks. Carbonate sedimentary rocks are mainly located in eastern Yunnan and northern Qinghai-Tibet Plateau. Acid plutonic rocks are typically distributed in the mountains surrounding the inland northeast, namely the Daxinganling Mountain and the hills in southern Guangdong and southwestern Guangxi. They are also distributed along the Brahmiputra river in the south part of the Qinghai-Tibet Plateau. The distribution of Acid plutonic

rocks is relatively scattered; there are many isolated Acid plutonic rocks distributions in different locations of contiguous China, accompanied by medium permeability and high porosity.

In summary, the The types of rocks in contiguous China are dominated by unconsolidated sediments and mixed sedimentary rocks. In 33.86% of the catchments, the dominant rock types occupy less than 50% of the catchment areas, and only 16.8% of basins are having a dominant rock type with an area fraction greater than 90%. Amongst 48754911 basins, 9.4% of basins have prevalent rock types wholly occupying the area.

5 Landcover





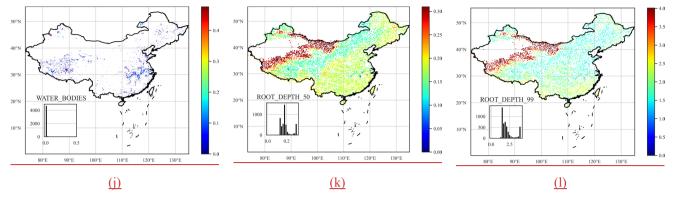


Figure 5. Maps: Distributions of land cover characteristics over contiguous China. The histograms indicate For land cover types, the number plot size is scaled by the size of catchments (out of 4875) in each bin, the land cover proportion.

We selected two indicators to characterize vegetation density and growth on the surface: Normalized difference vegetation index (NDVI) and Leaf area index (LAI). NDVI is an indicator with a valid range of -0.2 to 1, assessing whether the area being observed contains live green vegetation or the plants' health. However, NDVI is just a qualitative measurement of the vegetation density; it cannot provide a quantitative estimate of the vegetation density in the area. Moreover, NDVI often provides inaccurate vegetation density measurements, and only long-term measurement and comparison can ensure its accuracy. NDVI alone is not enough to estimate the state of plants in an area. Therefore, we have selected another indicator, LAI, to supplement the deficiencies of NDVI.

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LAI is defined as the total needle surface area per unit ground area and half of the entire needle surface area per unit ground surface area. It is a quantifiable value. It is functionally related to many hydrological processes like water interception (van Wijk and Williams 2005). (Buermann, Dong et al. 2001) verifies the validity of LAI used to characterize vegetation growth. The data sources used are The Terra Moderate Resolution Imaging Spectroradiometer (MODIS) Vegetation Indices (Didan 2015) for NDVI and Moderate Resolution Imaging Spectroradiometer (MODIS) (Myneni, Knyazikhin et al. 2015) for LAI. Followed (Addor, Newman et al. 2017), we determined maximum monthly LAI as an indicator characterising vegetation interception capacity and the maximum evaporative capacity and the difference between the maximum and minimum monthly LAI representing LAI's temporal variations.

Land cover classification refers to segmenting the ground into different categories based on remote sensing images. The Terra and Aqua combined Moderate Resolution Imaging Spectroradiometer (MODIS) Land Cover Type provides different results depending on the classification system used. Annual International Geosphere-Biosphere Programme (IGBP) classification is used for building the dataset, which is derived by the c4.5 decision tree algorithm. The IGBP classification system was formulated by the IGBP Land Cover Working Group in 1995, resulting in 17 categories of land cover types (Belward, Estes et al. 1999). (Friedl, Sulla-Menashe et al. 2010) compared the IGBP data of MODIS with other reference datasets and concluded that the MODIS

classification of IGBP has an accuracy of 75%. We determined the fraction of each land cover class for each basin based on the Terra and Aqua combined Moderate Resolution Imaging Spectroradiometer (MODIS) Land Cover Type (Sulla-Menashe and Friedl 2018), which differentiates our dataset from CAMELS and CAMELS-CL (only calculated the proportion of the dominant types).

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Followed (Addor, Newman et al. 2017), we also computed the average rooting depth (50% and 90%) for each catchment based on the IGBP classification using a two-parameter method (Zeng 2001). The root depth distribution of vegetation affects the ground's water holding capacity and the topsoil layer's annual evapotranspiration (Desborough 1997). Many models use root depth as an essential parameter to characterize soil moisture absorption capacity. (Zeng 2001) developed a two-parameter asymptotic equation for estimating root depth distribution; the root depth distribution is global, derived based on the IGBP classification avoiding the problem of significantly different root distributions in various research. Figure 5(g) shows root depth distributions of different vegetation types, based on (Zeng 2001) is method. The 90% root depth is usually considered to be "rooting depth", among the 17 categories of IGBP, cropland has the smallest rooting depth, and open shrubland has the largest. The 90% root depth of all vegetation is less than 2 meters. The national distribution of catchments soil characteristics is shown in Fig. 5.

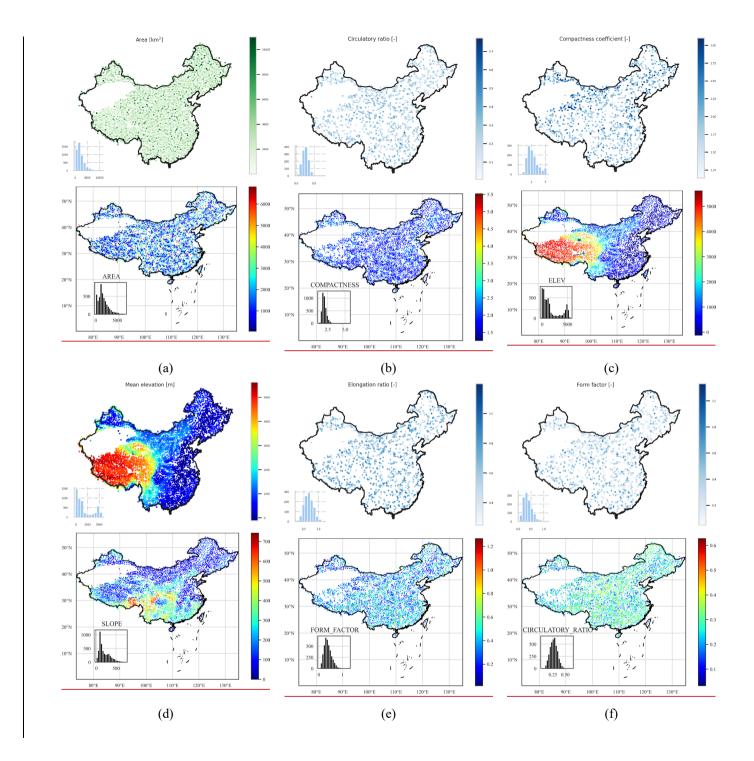
6 Location and topography

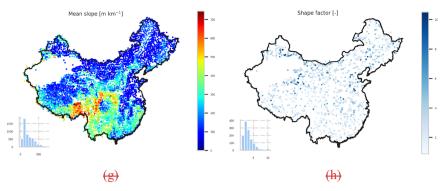
The catchments' boundary files are obtained from the global drainage basin dataset (Masutomi, Inui et al. 2009). The PDBDGDBD dataset was derived from digital elevation models (DEMs) with a high-resolution (100m-1km), and the errors were corrected by either automatic methods or manually. Additionally, PDBDGDBD also provides population and population density estimates for catchments, and these two indicators are also included in our dataset as a measure of human intervention. Global RunoffStreamflow Data Centre (Center 2005) discharge gauging stations were used for referencing the derived basins. In contiguous China, PDBDGDBD has- a high average match area rate (AMAR) and good geographic agreement with existing global drainage basin data in China. Based on the high-quality dataset, precise geographic and topographic information can be derived. See Fig. 6 for a summary.

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The topography attributes of each catchment are determined based on the ASTGTM product retrieved from https://lpdaac.usgs.gov, maintained by the NASA EOSDIS Land Processes Distributed Active Archive Center (LP DAAC) at the USGS Earth Resources Observation and Science (EROS) Center.





415 Figure 6. Maps Distributions of topographic characteristics over contiguous China. The histograms indicate the number of catchments (out of 4875) in each bin.

The physical characteristics of a catchment can affect the runoffstreamflow volume and the runoffstreamflow hydrograph of the catchment under a storm. To provide a complete description of the catchment shape, we computed several geometrical parameters of the catchment related to the runoffstreamflow process, (Fig. 6), including catchment form factor, shape factor, compactness coefficient, circulatory ratio and the elongation ratio (Subramanya 2013). A summary of the location and topography attributes can be found in Table 3.

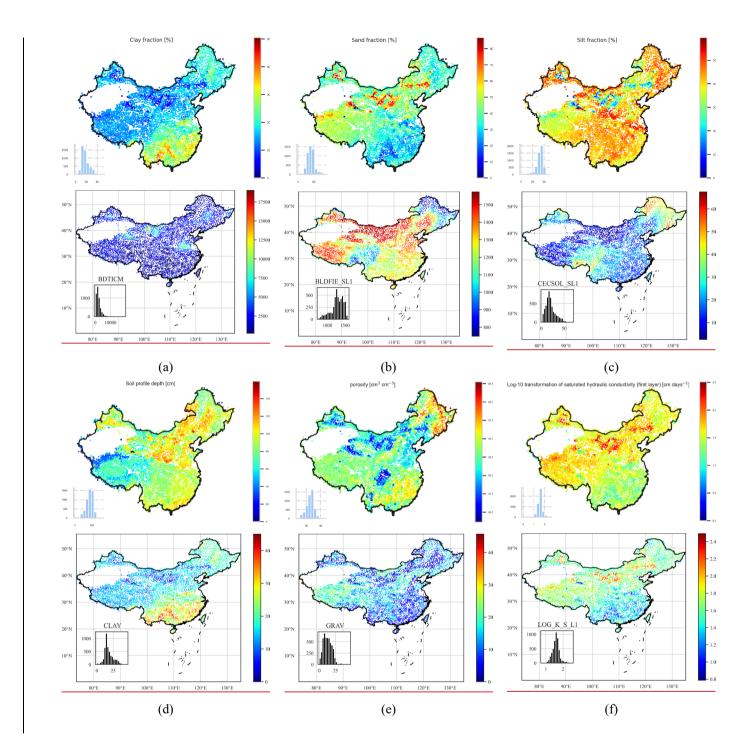
7 Soil

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The proposed dataset has a total of 54 soil attributes (Table 3) derived from (Hengl, Mendes de Jesus et al. 2017), (Dai, Xin et al. 2019) and (Shangguan, Dai et al. 2013). The summary result is shown in Fig. 7. Five categories of soil characteristics (pH in H2O, organic carbon content, depth to bedrock, cation-exchange capacity, and bulk density) are determined from SoilGrids. SoilGrids (Hengl, Mendes de Jesus et al. 2017) provides global predictions for soil properties including organic carbon, bulk density, cation exchange capacity (CEC), pH, soil texture fractions and coarse fragments by fusing multiple data sources including MODIS land products, SRTM DEM, climatic images and global landform and lithology maps at the 250m resolution-(Fig. 7). SoilGrids made predictions based on machine learning algorithms and many covariates ovariates layers primarily derived from remote sensing data. SoilGrids has soil characteristics for several soil depths.



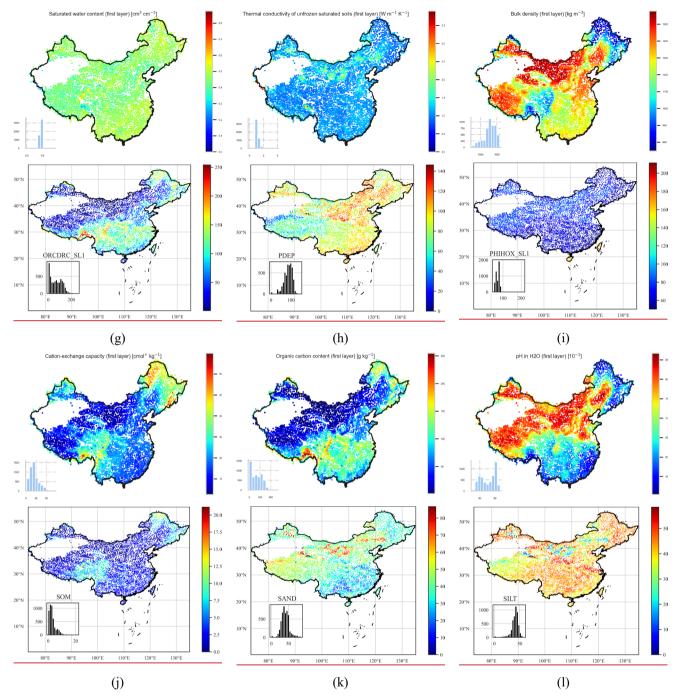


Figure 7. <u>Maps: Distributions</u> of soil characteristics over contiguous China. The histograms indicate the number of catchments (out of 4875) in each bin.

435 Unlike Different from CAMELS, whose reported results are obtained by a linear weighted combination of the different soil layers, and CAMELS-BR, whose products are soil characteristics at a depth of 30cm. We computed soil characteristics at all soil layers provided by Soil Grids such that advanced models can learn directly from the raw inputs Soil Grids 250m.

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To be consistent with CAMELS, we also We determined saturated water content and saturated hydraulic conductivity (Dai, Xin et al. 2019). We also introduced thermal conductivity of unfrozen saturated soils (Dai, Xin et al. 2019). (Dai, Xin et al. 2019) Based on the same dataset, we also introduced the thermal conductivity of unfrozen saturated soils. Dai, Xin et al. (2019) provides a global estimation of soil hydraulic and thermal parameters using multiple Pedotransfer Functions (PTFs) based on SoilGrids: the SoilGrids250m dataset. Based on the SoilGridsSoilGrids250m and GSDE (Shangguan, Dai et al. 2014) datasets, (Dai, Xin et al. 2019) produced six soil layers with a spatial resolution Dai, Xin et al. (2019) produced six soil layers with a spatial resolution of (Dai, Xin et al. 2019) is the same as the SoilGrids250m, with six intervals of 0–0.05 m, 0.05–0.15 m, 0.15–0.30 m, 0.30–0.60 m, 0.60–1.00 m, and 1.00–2.00 m. Same as the methods applied to SoilGrids, we determined We determine and records catchment soil characteristics for all these layers.

To provide even more complete description of the soil In addition, we determined seven more soil characteristics (Shangguan, Dai et al. 2013) including soil profile depth, porosity, clay/silt/sand content, rock fragment, and soil organic carbon content. (Shangguan, Dai et al. 2013) Shangguan, Dai et al. (2013) provides physical and chemical attributes of soils derived from 8979 soil profiles at 30×30 arc-second resolution, the polygon linkage method was used to derive the spatial distribution of soil properties. The profile attribute database and soil map are linked under a framework avoiding uncertainty in taxon referencing.

Depth to bedrock controls many physical and chemical processes in soil. The distribution of depth to bedrock in contiguous China is characterised by (i) low in the mountainous areas, such as Yunnan province and Chongqing City; (ii) high in barren areas, e.g. North and Northwest China. The introduced soil pH value is crucial since it influences many other physical and chemical soil characteristics. The spatial variability of soil pH in contiguous China is characterised by (i) soils in southern contiguous China are acid to strongly acid; (ii) soils in northern China are natural or alkaline; (iii) soils in northeastern forested areas are also acid (pH < 7.2). Cation exchange capacity can be seen as a measure of soil fertility since it measures how much nutrient the soil can store such that it influences the growth of the vegetations vegetation. Cation exchange capacity is positively correlated with soil organic matter content and clay content, which Cation exchange capacity is generally low in sandy and silty soils. The spatial variability of Cation exchange capacity in contiguous China is characterised by (i) high in peat and forested areas in Qinghai-Tibet Plateau, central and northeast China (ii) The Cation exchange capacity in the desert area such as the northwest is extremely low. Soil hydraulic and thermal properties are greatly affected by soil organic matter (SOM). Soil organic matter has a similar distribution to the cation exchange capacity: high in the peat and forested areas such as northeast China and low in the north and northwest.

8 Meteorological time series

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Table 4: Summary table of catchment meteorological time series available in the proposed dataset

| Variable | Description | Unit |
|----------|--|--------------------|
| prs | catchment daily averaged ground pressure | hPa |
| tem | catchment daily averaged temperature at 2 m above ground | °C |
| rhu | catchment daily averaged relative humidity | - |
| pre | catchment daily averaged precipitation | mm d ⁻¹ |
| evp | catchment daily averaged evaporation measured by ground instruments | mm d ⁻¹ |
| win | catchment daily averaged wind speed at 2 m above ground | m s ⁻¹ |
| ssd | catchment daily averaged sunshine duration | h d ⁻¹ |
| gst | catchment daily averaged ground surface temperature | °C |
| pet | catchment daily averaged potential evapotranspiration determined by Penman's equation (see | mm d ⁻¹ |
| | Appendix A) | |

There have been many studies based on SURF_CLI_CHN_MUL_DAY in China (Liu, Xu et al. 2004, Xu, Gao et al. 2009, Huang, Han et al. 2016, Liu, Zheng et al. 2017), such as trend analysis of the pan evaporation (Liu, Yang et al. 2010). Still, there has not yet been a large-scale basin-oriented meteorological time series dataset in contiguous—China. Researchers still need to do repeated works to extract historical meteorological data from the SURF_CLI_CHN_MUL_DAY dataset for the research. For the first time, we release a catchment_scale meteorological time series dataset. We will also The open-source the sourced code for researchers to can generate any catchment's meteorological time series within contiguous—China. The basin-oriented dataset provides meteorological time series for 48754911 basins from 1990 to 20182020 based on the China Meteorological Data Network. Meteorological time series includes pressure, temperature, relative humidity, precipitation, evaporation, wind speed, sunshine duration, ground surface temperature and potential evapotranspiration (see __(Table 4 for a summary).).

The meteorological time series data from 1951 to 2010 is derived based on the "1951-2010 China National Ground Station Data Corrected Monthly Data File Basic Data Collection" data construction project. Other data include monthly reported data to the National Meteorological Information Centre by the provinces, and hourly and daily data uploaded by automatic ground stations in real-time. The SURF_CLI_CHN_MUL_DAY dataset is quality controlled, the quality and completeness of each variable are significantly improved compared to the previous similar products. MDuring the development of the dataset, missing data were filled by interpolating its nearest stations.

Figure 2 presents the variation of the distribution number of the observation sites. The start date of the recording is 1951, but because the early site distribution is sparse, we only used records from 1990 to 20182020 to construct the dataset to ensure the data quality. The interpolation method used is the Inverse distance weighting since it shows better performance than other comparators. Catchment-scale raster is extracted from the interpolated national raster using the open-source rasterio package. For all variables, we take the arithmetic mean on the extracted catchment raster as the catchment mean. Potential interpolation methods. In addition, potential evapotranspiration (PET) is estimated based on Penman's Equation (Appendix A) and other catchment meteorological variables.

9 Normal Camels YR - Normalized Catchment attributes and meteorology for Yellow River basin

Apart from the dataset providing the catchment attributes and meteorological forcing for contiguous China, we also offer a self-contained dataset covering the Yellow River basin with normalized streamflow measurements. The streamflow data are normalized to have zero mean and a standard deviation of 1 for each basin. The Normal Camels YR dataset is designed to support machine learning and deep learning research related to hydrology. In particular, fifty four watersheds are less affected by human activities (selection is based on the Global Reservoirs and Dam databases (GRanD) (Lehner, Liermann et al. 2011) which provides the locations of reservoirs and dams globally), which makes them suitable for rainfall runoff modelling research. For most machine learning and deep learning algorithms, data normalization will not affect model performance (e.g., neural network based and tree based algorithms). Besides, other research, such as trend analysis, can also be carried out. The Normal Camels YR dataset is self-contained to fully describe the Yellow River basin and is particularly helpful for the hydrology research of the Yellow River.

During the dataset development, basins with too few observations are removed, resulting in discontinuous basin identifiers. Normal Camels YR covers 102 gauges in the Yellow River basin, providing basin boundary shapefiles, static attributes and normalized streamflow measurements for each basin. The covered basins have areas ranging from 134 to 804,421 square kilometres. The time resolution of streamflow measurements is seven days, and the mean length of records of the streamflow measurements is 684, which means the mean period of the streamflow measurements for each basin is over 13 years. Meteorological variables included in Normal-Camels-YR is slightly different; it introduced daily maximum and minimum for some variables (Table 5).

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⁶⁻https://github.com/mapbox/rasterio

9 HydroMLYR: Hydrology dataset for Machine Learning in YRB

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In addition to the basin-wise static attributes provided in CCAM, we propose HydroMLYR, a hydrology dataset for machine learning research in the YRB (Fig. 1). HydroMLYR includes standardized streamflow measurements for 102 basins. The streamflow data is seven-day averaged and standardized basin-wise to have zero mean and a standard deviation of 1 (Fig. 8). The HydroMLYR dataset is proposed to support machine learning or deep learning hydrology research (e.g., neural network-based and tree-based algorithms). It can be used in two cases: (1) to develop machine learning models on the YRB or (2) when it is desirable to verify the generalization ability of a machine learning model on YRB.

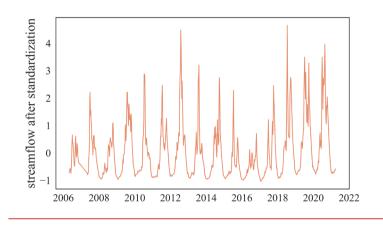


Figure 8: Examples of standardized runoff

The dataset provides 40 natural basins in the dataset which are not affected by reservoirs and dams. The selection is based on a newer version⁷ of the Global Reservoirs and Dam databases (Lehner, Liermann et al. 2011) which provides the locations of reservoirs and dams globally. HydroMLYR covers 102 basins in the YRB, including basin boundary shapefiles, static attributes, and standardized streamflow measurements for each basin. The covered basins have areas ranging from 134 to 804,421 square kilometres. Therefore, modelling on a large scale of the YRB is also possible. Meteorological records in HydroMLYR introduced daily maximum and minimum for some forcing variables (Table 5).

The original streamflow observations are not continuous. The average record length is 11.3 years. Although the development of machine learning models does not necessarily require the data to be continuous, we separately provide continuous streamflow observations with an average record length of 8.3 years.

Table 5: Meteorological variables provided in Normal Camels YR, the time series length is 22 years (1999-2020) HydroMLYR

| Attribute name | Description | Unit |
|----------------|---|------------------------------------|
| evp | catchment daily averaged evaporation (observations) | 0.1 -mm d ⁻¹ |
| gst_mean | catchment daily averaged ground surface temperature | 0.1 °C |

⁷ http://globaldamwatch.org/data/#core_global

| gst_min | catchment daily minimum ground surface temperature | 0.1 °C |
|----------|--|------------------------------------|
| gst_max | catchment daily maximum ground surface temperature | 0.1 °C |
| pre | catchment daily averaged precipitation | 0.1 -mm d ⁻¹ |
| prs_mean | catchment daily averaged ground surface pressure | 0.1 -hPa |
| prs_max | catchment daily maximum ground surface pressure | 0.1 -hPa |
| prs_min | catchment daily minimum ground surface pressure | 0.1 -hPa |
| rhu | catchment daily averaged relative humidity | - |
| ssd | catchment daily averaged sunshine duration | 0.1 h |
| tem_mean | catchment daily averaged temperature | 0.1 °C |
| tem_min | catchment daily minimum temperature | 0.1 °C |
| tem_max | catchment daily maximum temperature | 0.1 °C |
| win_max | catchment daily maximum wind speed | 0.1 -m s ⁻¹ |
| win_mean | catchment daily averaged wind speed | 0.1 -m s ⁻¹ |
| | | |

10 Data and code availability and software packages used.

The proposed dataset is freely available at http://doi.org/10.5281/zenodo.47040175137288. The files provided are (i) several separate files containing 120+ catchments attributes, (ii) the daily meteorological time series in a zip file, (iii) the catchment boundaries used to compute the attributes and extract the time series, (iv) the Normal Camels YRHydroMLYR dataset, (v) an attribute description file and (v) a readme file. The code used to generate the dataset is mainly based on several publicly available packages: rasterio, gdal 8, pyshp 9, geopandas 10, fiona 11, and xarray 12. Complement code for generating any watershed's dataset will be released soon.

11 Conclusion

The <u>CCAM</u> dataset proposed in this paper provides a novel dataset for hydrological research in contiguous China. In the study <u>China. area, there is no catchment attributes dataset has been proposed before, either a catchment scale time series</u>

⁸ https://github.com/OSGeo/gdal

⁹ https://github.com/GeospatialPython/pyshp

¹⁰ https://github.com/geopandas/geopandas

⁴¹ https://github.com/Toblerity/Fiona

¹² https://github.com/pydata/xarray

meteorological dataset. All catchments delaminated from the DEM are studied, covering contiguous entire China. The dataset includes daily meteorological forcing time-series data including precipitation, temperature, potential evapotranspiration, wind, ground surface temperature, pressure, humidity, sunshine duration and derived potential evapotranspiration of 48754911 catchments. The proposed time series dataset is derived based on the quality-controlled site observation dataset, SURF CLI CHN MUL DAY. We will also release the complement code for generating any shapefile's meteorological time series within contiguous China based on the SURF CLI CHN MUL DAY dataset (freely available for Chinese researchers). The dataset has longer time series (from 1990 to 2018) and more meteorological variables than the previously proposed datasets. The dataset also dataset. CCAM includes 120+ catchment attributes, including soil, land cover, geology, climate indices and topography for each catchment. -We produced a series of maps depicting the catchment attributes distributions in contiguous China. These maps present regional changes of various features; we also describe estimate the relationships between them. The integration of based on Kendall's correlation. Integrating multiple data sources into one dataset at a catchment- scale dramatically simplifies the data compilation process in research. Based on the dataset, we CCAM can help test hypotheses and formulate valid conclusions under various conditions, not just limited to a few specific locations-Together with the Normal Camels YR dataset, the proposed dataset can and help explore how different basin characteristics influence hydrological behaviours, learn the migration of hydrological behaviours between different basins, and to-develop general frameworks for large-scale model evaluation and benchmarking in China. A limitation of the study is the lack of estimation of the uncertainty of the meteorological time series. An alternative is to evaluate the uncertainty of the basin-wise meteorological data based on multiple independent data sources, but there are few data that provide as many data types as SURF CLI CHN MUL DAY. Hence, it poses a challenge for evaluating the uncertainty of these eight meteorological variables, which is left for future studies.

Appendix A: Modified Penman's equation

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Penman's equation (Subramanya 2013), incorporating some modifications to the original formula, is:

$$PET = \frac{AH_n + E_a \gamma}{A + \gamma}$$

where PET is the daily potential evapotranspiration in mm per day; A is the slope of the saturation vapour pressure (ew) vs temperature (t) curve at the mean air temperature, in mm of mercury per Celsius; Hn is the net radiation in mm of evaporable water per day; Ea is a parameter including wind speed and saturation deficit; γ is the psychrometric constant = 0.49 mm of mercury per Celsius.

5 The relationship between ew and t is defined as:

$$e_w = 4.584 \exp\left(\frac{17.27t}{237.3 + t}\right)$$

The following equation estimates the net radiation:

$$H_n = H_a(1-r)\left(a + b\frac{n}{N}\right) - \sigma T_a^4 \left(0.56 - 0.092\sqrt{e_a}\right) \left(0.10 + 0.90\frac{n}{N}\right)$$

where Ha is the incident solar radiation outside the atmosphere on a horizontal surface, expressed in mm of evaporable water per day (a function of the latitude and period of the year as indicated in Table A1); a is a constant depending upon the latitude ϕ and is given by $a = 0.29 \cos \phi$; b is a constant = 0.52; n is the sunshine duration in hours; N is the maximum possible hours of bright sunshine (a function of latitude, see Table A2); r is the reflection coefficient; σ is the Stefan-Boltzman constant = 2.01×10^{-9} mm/day; Ta is the mean air temperature in degrees kelvin; ea is the actual mean vapour pressure in the air in mm of mercury.

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Table A1:: Mean Monthly Solar Radiation, Ha in mm of Evaporable Water/Day

| North latitude | Jan | Feb | Mar | Apr | May | Jun | Jul | Aug | Sep | Oct | Nov | Dec |
|----------------|------|------|------|------|------|------|------|------|------|------|------|------|
| 0° | 14.5 | 15.0 | 15.2 | 14.7 | 13.9 | 13.4 | 13.5 | 14.2 | 14.9 | 15.0 | 14.6 | 14.3 |
| 10° | 12.8 | 13.9 | 14.8 | 15.2 | 15.0 | 14.8 | 14.8 | 15.0 | 14.9 | 14.1 | 13.1 | 12.4 |
| 20° | 10.8 | 12.3 | 13.9 | 15.2 | 15.7 | 15.8 | 15.7 | 15.3 | 14.4 | 12.9 | 11.2 | 10.3 |
| 30° | 8.5 | 10.5 | 12.7 | 14.8 | 16.0 | 16.5 | 16.2 | 15.3 | 13.5 | 11.3 | 9.1 | 7.9 |
| 40° | 6.0 | 8.3 | 11.0 | 13.9 | 15.9 | 16.7 | 16.3 | 14.8 | 12.2 | 9.3 | 6.7 | 5.4 |
| 50° | 3.6 | 5.9 | 9.1 | 12.7 | 15.4 | 16.7 | 16.1 | 13.9 | 10.5 | 7.1 | 4.3 | 3.0 |

The parameter Ea is estimated as:

$$E_a = 0.35 \left(1 + \frac{u_2}{160}\right) (e_w - e_a)$$

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where u2 is the wind speed at 2m above ground in km/day; ew is the saturation vapour pressure at mean air temperature in mm of mercury; ea is the actual vapour pressure.

Table A2-: Mean Monthly Values of Possible Sunshine Hours, N

| North latitude | Jan | Feb | Mar | Apr | May | Jun | Jul | Aug | Sep | Oct | Nov | Dec |
|----------------|------|------|------|------|------|------|------|------|------|------|------|------|
| 0° | 12.1 | 12.1 | 12.1 | 12.1 | 12.1 | 12.1 | 12.1 | 12.1 | 12.1 | 12.1 | 12.1 | 12.1 |
| 10° | 11.6 | 11.8 | 12.1 | 12.4 | 12.6 | 12.7 | 12.6 | 12.4 | 12.9 | 11.9 | 11.7 | 11.5 |
| 20° | 11.1 | 11.5 | 12.0 | 12.6 | 13.1 | 13.3 | 13.2 | 12.8 | 12.3 | 11.7 | 11.2 | 10.9 |
| 30° | 10.4 | 11.1 | 12.0 | 12.9 | 13.7 | 14.1 | 13.9 | 13.2 | 12.4 | 11.5 | 10.6 | 10.2 |
| 40° | 9.6 | 10.7 | 11.9 | 13.2 | 14.4 | 15.0 | 14.7 | 13.8 | 12.5 | 11.2 | 10.0 | 9.4 |
| 50° | 8.6 | 10.1 | 11.8 | 13.8 | 15.4 | 16.4 | 16.0 | 14.5 | 12.7 | 10.8 | 9.1 | 8.1 |

Appendix B: Correlation analysis of catchment attributes

- To explore the potential connections between various types of watershed attributes, we did correlation analysis using the Kendall rank correlation coefficient (Kendall 1938). Kendall rank correlation coefficient is a measure of rank correlation: the similarity of the sort order of the two sets of data. Kendall correlation will be high if the orderings of the observations of two variables are similar. Kendall correlation avoids the assumption of linear relationship and that the distribution should be normal and continuous (e.g., Pearson correlation coefficient; the results can be found in). When the relationship is not exactly linear, using Pearson correlation will miss out on information that Kendall could capture. Table B1, which shows the top five most relevant attributes for each attribute, and the Fig. S1, the correlation matrix. The analysis result shows that the correlations between variables are consistent in line with general understanding, justifying the rationality of the dataset, to name a few:
 - (1) Subsurface permeability and porosity are highlymost correlated with geological attributes.
 - (2) LAI and NDVI have a high positive correlation (0.866).
- 605 (3)(2) Root depth is are most positively correlated with each other but most negatively correlated with the fraction of barren land cover types.
 - (3) Urban and built ups are most positively correlated with population density.
 - (4) In China, the savanna is mainly distributed in the southern coastal areas, resulting in that it is <u>most</u> positively correlated with <u>average rainfall (0.604)-mean precipitation.</u>
- 610 (5) Sand is <u>most</u> positively correlated with <u>the saturated hydraulic conductivity (0.86)</u> while the clay is <u>strongly</u> negatively correlated (-0.763), and <u>catchments</u> with <u>a lot of rainfall are less likely to have soil with high hydraulic conductivity (-0.647).</u>
 - (6) High altitude catchments tend to have lower saturated water content (-0.705).

Table B1: The top five most relevant characteristics for each attribute (different soil layers for the same attribute are excluded, e.g., phihox_sl2 is not included in the top five most relevant attributes of phihox_sl1 though they are highly correlated)

| Attribute | 1 st | 2 nd | 3^{rd} | 4 th | 5 th |
|---------------|---------------------------------------|--------------------------------------|--|---|---------------------------|
| high_prec_fre | low_prec_durroot_de | root_depth_50(- | root_depth_99(- | barren(| pet_meantksatu_11(- |
| q | <u>pth_50(</u> -0. 58 196) | <u>grassland(</u> 0.438 <u>175</u>) | 0.4 <u>36</u> 171) | som(0.39136) | 0. 261 133) |
| high_prec_du | elev(theta_s_l6(- | theta_s_ 16 <u>15</u> (- | p_seasonality(prs | theta_s_l5(- | rhu_meantheta_s_14 |
| r | 0.544 <u>277</u>) | 0. 503 <u>234</u>) | _mean (- 0.49 <u>233</u>) | <u>elev(</u> 0.458 <u>211</u>) | (-0.431 <u>201</u>) |
| low_prec_fre | pre_mean(- | ssd_meanaridity(0.84 | phihox_sl7ssd_mea | phihox_sl6(rhu_mea | phihox_s15s17(0.81 |
| q | 0. 881<u>766</u>) | ± <u>745</u>) | <u>n</u> (0. 825 <u>652</u>) | <u>n(-</u> 0. 818 <u>627</u>) | 4 <u>588</u>) |
| low proc dur | barrenaridity(0.7287 | rhupre_mean(- | evpssd_mean(0.72 | ndvirhu_mean(- | phihox_sl7reet_d |
| low_prec_dur | <u>8</u>) | 0. 723 <u>768</u>) | 4 <u>731</u>) | 0. 68 4 <u>709</u>) | epth_99(0.66 <u>579</u>) |

| frac_snow_da ily | temgst_mean(- 0.951802) | gsttem_mean(- 0.949 <u>792</u>) | ssd_mean_lat(0.7775 75) | <u>broadleaf_tree(-</u> 0.762512) | n_min(pre_mean(- 0.703436) |
|---------------------------|--|--|---|---|--|
| prs mean _{p_sea} | pre_mean(elev(- 0.901678) | rhu_meanlon(0.765 <u>55</u> 2) | ssd <u>rhu</u> _mean(- (0.764432) | low_prec_freq(0.712)urban_and_b uilt- up_land(0.427) | frac_snow_dailybarre n(-0.68341) |
| petpre_mean | cecsol_sl2aridity(- 0.66913) | <u>dur(-0.634768)</u> | cecsol_sl3low_prec _freq(-0.628766) | gstssd_mean((- 0.622723) | bldfie_sl1 <u>rhu_mean</u> (0.608 <u>712</u>) |
| preevp_mean | p_seasonalityaridity(0 .901643) | low_prec_freqndvi_m ean(-0.881632) | ssdrhu_mean(- 0.858 <u>617</u>) | rhussd_mean(0.832 598) | phihox_sl7 <u>lai_dif</u> (- 0. <u>819593</u>) |
| temgst_mean | gsttem_mean(0.9929 24) | frac_snow_daily(- 0.951802) | pre_mean(lat(- 0.747 <u>512</u>) | oadleaf_tree(0.50 7) | pet_meanp_seaso nality(0.681442) |
| prs <u>rhu</u> mean | elevaridity(- 0.889 <u>751</u>) | e_max(ssd_mean(- 0.707746) | lonpre_mean(0.70 7712) | <u>e_min(low_prec_d</u> <u>ur(-0.707709)</u> | rhu_mean(low_prec_freq(-0.603627) |
| rhupet_mean | -0.887451) | pregst_mean(0.8324 42) | evp_meancecsol_sl 3(-0.823441) | ndvi_mean(cecsol_s l1(-0.813422) | low_prec_freqcecsol _s14(-0.80342) |
| evpssd_mean | ndvi_mean(aridity(0.845753) | rhu_mean(- 0.823746) | ssd_meanlow_prec_dur(0.756731) | e_minpre_mean(- 0.731723) | low_prec_freq(0.7 3652) |
| win_mean | ssd_mean(0. 581 <u>426</u>) | frae_snew_daily(wood y_savanna(- 0.571393) | tem_mean(- 0.52 <u>379</u>) | gst_mean(- 0. 507 <u>377</u>) | low_pree_freq(mixed _forest(_0.477363) |
| ssdtem_mean | rhugst_mean(- (0.887924) | pre_meanfrac_snow_daily(-0.858792) | low_pree_freqeverg reen_broadleaf_t ree(0.841493) | frae_snow_dailypop_dnsty(0.777475) | p_seasonality <u>lat(</u> - 0.764 <u>474</u>) |
| | | | | | |
| p seasonality gst_mean | temrhu_mean((- 0.992421) | frae_snow_dailytem_ mean(-0.949397) | pregst_mean((- 0.743393) | n_min(ssd_mean(0.69339 3) | low_prec_dur(0.69 3375) |
| | | _ | | ssd_mean(0.69339 | low_prec_dur(0.69 |

| slopegeel_por | su(<u>lat(</u> -0.627 <u>374</u>) | pa <u>bdticm</u> (- 0.575 <u>348</u>) | phihox_sl1(win_me an(-0.46341) | phihox_sl3mixed_fo rest(0.454341) | n_needleleaf_tree(0.453327) |
|----------------------------|---|--|--|---|--|
| ig lon | now and ice(clev(- 0.471 <u>585</u>) | tksatu_l5prs_mean(0. 324552) | tksatu_13(evp_mea n(-0.3185) | tksatu_14(barren(- 0.306482) | tksatu_l2ndvi_mean (0.27547) |
| <u>elev_{pa}</u> | geol_porosityprs_mea n(-0.575678) | phihox_sl1]on(- 0.314 <u>585</u>) | nd_built- up_land(- 0.302485) | phihox_sl2pop_dnst y(-0.301481) | phihox_sl4cropland(-0.297456) |
| se <u>lat</u> | geol_porosity(frac_snow_daily(0. 362575) | evergreen_broadle af_tree(-0.548) | n_maxgst_mean(- 0.317 <u>512</u>) | lattem_mean(- 0.317474) | n_min(low_prec_freq(0.3 16437) |
| su pop | geel_porosityurban_a nd_built- up_land(0.627618) | bdtiemcropland(0.59 9519) | cropland(aridity(- 0.468 <u>511</u>) | <u>phihox_sl1pre_mea</u> <u>n(0.44505)</u> | phihox_sl4rhu_mea n(0.439492) |
| pop dnstysm | geol_permeability(- 0.403)urban_and_bu ilt-up_land(0.639) | suaridity(-0.385 <u>538</u>) | cropland← (0. 268 533) | bdtiem(- pre_mean(0.23353 3) | e_maxssd_mean(- 0.228521) |
| <u>length</u> vi | deciduous broadleaf treearca(0.214684) | geol_porosityform_fa ctor(-0.18398) | lei_mexshape_fact or(0.165398) | lai_dif(elongation_r atio(-0.459398) | e_max_compactness coefficient(0.157 363) |
| <u>areamt</u> | geol_porosity(length(0.412684) | evergreen needleleaf treepop(0.32723) | oredre_sl3pa(0.265 194) | oredre_sl4(circulato ry_ratio(- 0.258187) | bldfie_sl5(0.254)compactness _coefficient(0.187 |
| ∞form factor | geol_permeability(0.408)elongation_rat io(1.0) | su(_shape_factor(- 1.0.287) | circulatory_ratio(0.206435) | mpactness_coeffic ient(-0.435) | tksatu_16length(- 0.156398) |
| pishape_factor | deciduous broadleaf tree(0.299)clongation _ratio(-1.0) | geol_porosity(0.208)form_factor(- 1.0) | e_max(circulatory _ratio(_0.161435) | loncompactness_c oefficient(0.16143 5) | e_minlength(0.1639 <u>8</u>) |
| vacompactnes s_coefficient | geol_porosity(- 0.218)circulatory_ra tio(-1.0) | high_pree_dur(elongat ion_ratio(- 0.191435) | shape_factor(0.16 7435) | gst_meanform_fact or(-0.16435) | eu(- length(0.16363) |

| circulatory ra tiowb | water bodies(0.674)compact ness_coefficient(- 1.0) | permanent wetlandelongation_r atio(0.379435) | <u>shape_factorroot</u> <u>_depth_50(</u> - 0.164 <u>435</u>) | theta_s_13_form_fact or(0.148435) | theta_s_l4(length(- 0.147363) |
|--|--|---|--|--|---|
| pbelongation ratio | theta_s_16(shape_factor(- 1.0.137) | theta_s_15(- form_factor(1.0.133) | elev(m)(circulatory ratio(0.124435) | theta_s_14compactn ess_coefficient(- 0.114435) | prs_meanlength(- 0.102398) |
| <u>lai dif</u> vb | cecsol_sl2ndvi_mean (0.222808) | cecsol_sl3(barren(- 0.213642) | cecsol_sl1(aridity(- 0.212638) | (0.211609) | vanna (0.208607) |
| ndlai_max | icendvi mean (0.206 | theta_s_12barren(- 0.154 <u>614</u>) | theta_s_l3aridity(- 0.151613) | theta_s_II(woody_savanna(0 .144612) | tksatu_14(phihox_sl 2(-0.136602) |
| ndvi_mean _{py} | phihox_sl1(lai_dif(0.214808) | phihox_sl2(lai_max(0.207779) | phihox_sl3barren(- 0.207677) | phihox_sl4evp_mea n(-0.205632) | phihox_sl5aridity(- 0.202607) |
| root depth 5 | tksatu_13(grassland(- 0.07485) | tksatu_14pet_mean(0. | barren(0.064 <u>212</u>) | high_prec_freq(- tksatu_12(0.061196) | pdep(- tksatu_11(0.061 <u>176</u> |
| root_depth_9 | ndvi_mean(grassland | phihox_sl4(| phihox_sl2cropland | phihox_sl5pdep(- | phihox_sl6lon(- |
| <u>9lai_dif</u> | <u>(-0.866339</u>) | <u>barren(</u> 0.809337) | (-0. 807<u>336</u>) | 0. 807 <u>284</u>) | 0. 807 283) |
| lai_maxevergre en_needleleaf tree | ndvi_meanmixed_for est(0.856572) | phihox_sl4(woody_savanna(0.8 15481) | phihox_s15s17(-0.814416) | phihox_sl6(- 0. <u>814411</u>) | phihox_ <u>\$12\$15</u> (- 0. <u>\$13409</u>) |
| ndvi_meaneverg reen_broadlea f_tree | lai_dif(lat(-0.866548) | lai_max(phihox_sl7(- 0.856 <u>538</u>) | evp_meanphihox_s l6(-0.845529) | rhu_mean(phihox_sl 5(-0.813522) | barren(pre_mean(0.77251 2) |
| root_depth_50de ciduous_needl eleaf_tree | barrencecsol_sl1(0.8 56274) | low_prec_dur(bldfie_s 11(-0.626274) | grassland(cecsol_sl2(0.5372 72) | orcdrc_sl2(ndvi_ mean(-0.51327) | evp_meancecsol_sl 3(0.497262) |
| root_depth_99de ciduous_broa dleaf_tree | barrenmixed_forest(0.897604) | low_pree_durWoody savanna(0.66568) | ndvi_mean← (0. 628 524) | evp_meanlai_max(0 .604 <u>5</u>) | rhu_mean(- lai_dif(0.486497) |
| evergreen needleleaf | woody_savannasle pe(0.398713) | bldfie_sl4(- 0.391)deciduous_bro adleaf_tree(0.604) | bldfie_sl5(- 0.384)evergreen_n | bldfie_sl3phihox_sl 7(-0.372565) | bldfie_sl7phihox_sl 6(-0.366563) |

| treemixed_fore st | | | eedleleaf_tree(0. 572) | | • |
|---|---|---|---|---|---|
| evergreen broadleaf treeclosed_shr ubland | broadleaf tree(0.50 4217) | lai_maxsavanna(0.48 316) | phihox_sl7(- mixed_forest(0.4 77158) | lai_dif(tksatu_14(- 0.471 <u>153</u>) | phihox_sl6theta_s_l 2(-0.47142) |
| deciduous needleleaf treeopen_shru bland | woody savannahigh prec_d ur(0.241179) | eecsol_sl2(rhu_mean(-0.231174) | oredre_sl2elev(0.22 617) | petssd_mean(- (0.21517) | bldfie_sl1prs_mean(-0.214 <u>165</u>) |
| deciduous broadleaf treewoody_sav anna | lai_maxmixed_forest (0.459713) | lai_dif(phihox_sl7(- 0.452628) | <u>cecsol_sl1(phihox</u> <u>sl4(-</u> 0.433628) | bldfie_sl1phihox_sl 3(-0.413627) | e_max(phihox_sl6(-0.361627) |
| mixed forestsavanna | oredre_sllpre_mean(0.501606) | ural vegetaion (0.47 | hai_difwoody_sava nna(0.466604) | phihox_sl6aridity(- 0.462602) | phihox_sl7ssd_mea n(-0.461591) |
| closed shrublandgrassl and | theta_s_Hroot_depth _50(-0.084485) | natural_vegetaion(-0.363) | se(tem_mean(- 0.075344) | theta_s_l2gst_mean(-0.072344) | urban and built up land(0.064)root dep th 99(-0.339) |
| open shrublandperma nent_wetland | high_pree_durwater_b odies(0.155469) | theta_s_l6(savanna(0.151363) | rhu_mean(0.149)urban_and_ built- up_land(0.347) | prspre_mean(- (0.147343) | evp_meanpop(0.139 343) |
| woody savanna croplan d | uilt- up_land(0.633546) | lai_difpop_dnsty(0.6 31533) | phihox_s17(pop(0.592519) | phihox_sl6elev(- 0.59456) | lon(phihox_sl5(0.585417) |
| savannaurban_a nd_built- up_land | pre_meanpop_dnsty(0.604639) | phihox_sl7(- pop(0.55618) | claycropland (0.54 7 <u>546</u>) | phihox_sl6elev(- 0.543485) | phihox_sl5(- 0.537)cropland_nat ural_vegetaion(0. 428) |
| nd_natural_ve getaion | root_depth_50(savanna(0.537605) | tem <u>rhu</u> _mean(- (0.496 <u>546</u>) | gst_meanaridity(- 0.491 <u>523</u>) | frae_snow_daily(ssd_ mean(-0.46952) | phihox_sl6pre_mea n(0.43851) |

| permanent wetlandsnow_a nd_ice | wbig(0. 37 9 <u>431</u>) | water bodiesbarren(0.34937 9) | p_seasonality(lon(- 0.3373) | pre_meanelev(0.248 369) | pop_dnsty(pdep(- 0.23354) |
|---|---|--|---|--|--|
| eroplandbarren | su(ndvi_mean(- 0.468677) | lon(lai_dif(- 0.412642) | e_min(lai_max(- 0.412614) | e_maxaridity(0.412 581) | evp_mean(elev(0.388574) |
| urban and built up landwater_bod ies | pop_dnstypermanent _wetland(0.811469) | рор wb(0.399 <u>39</u>) | nd natural veget aion(0.28617) | d built- up_land(0.261158) | elev(-0.244 <u>154</u>) |
| geol permeab ilityeropland/nat ural vegetaion | ssd_meansm(- 0.458345) | savanna <u>Su</u> (0. 381 <u>326</u>) | rhu_mean(<u>SS(</u> - 0.371 <u>316</u>) | frac_snow_daily(bdticm(0.367228) | tem_meanpdep(0.36 4 <u>161</u>) |
| geol_porositys | tksatu_15su(0.568 <u>455</u> | tksatu_13(pa(- 0.561417) | tksatu_14(woody_s avanna(- 0.533323) | tksatu_12phihox_s13 (0.506315) | tksatu_11phihox_sl4 (0.503314) |
| barren ig | root_depth_99snow_a nd_ice(0.897431) | root_depth_50elev(0.8 56194) | ndvi_meantheta_s_ 12(-0.772185) | low_prec_dur(pdep(- 0.728184) | evp_mean(theta_s_1 3(-0.698182) |
| water bodies pa | wb(geol_porosity(- 0.674 <u>417</u>) | $\frac{\text{permanent}}{\text{wetland}\underline{\text{mt}}}(0.349\underline{3})$ | root_depth_50(- pi(0.192295) | root_depth_99(va(0.154271) | theta_s_l3vi(0.153 <u>24</u> 6) |
| <u>sc</u> length | ####@col_porosity(- 0.849285) | circulatory_ratio <u>lat(</u> -0.491 <u>264</u>) | elongation_ratio_bdti cm(-0.45126) | form_factor(slope(0.436246) | eompactness_coefficie ntmixed_forest(0.2 92231) |
| area S <u>U</u> | length bdticm(0.8495) 2) | popgeol_porosity(0. 418 <u>455</u>) | eireulatory_ratioW00 dy_savanna(- 0.255349) | eecsol_sllgeol_per meability(0.14232 6) | bldfie_sl2(- phihox_sl7(0.1383 26) |
| form_factorSM | elongation_ratio(geol_permeability(-0.992345) | circulatory_ratio(su(- 0.647283) | shape_factorbdticm (-0.506228) | lengthcropland(- 0.436 <u>199</u>) | eompactness_coefficie nt(- 0.383)clev(0.194) |
| shape_factorVi | compactness_coefficien tpa(0.786246) | elongation_ratio(- pi(0.566203) | form_factor(- va(0.506171) | eirculatory_ratiogeol _porosity(- 0.372169) | lengthdeciduous br oadleaf_tree(0.266 166) |
| compactness_coe fficientmt | shape_factorpa(0.7863 | geol_porosityeireul atory_ratio(- 0.594286) | elongation_ratio(pi(0.42+199) | form_factor(0.383)deciduous_br oadleaf_tree(0.18 7) | l ength area(0. 292 18) |

| circulatory_ratioS S elongation_ratiop | elongation_ratio(geol_permeability(- 0.651316) form_factorpa(0.9922 | form_factor(Su(- 0.647 <u>17</u>) | eompactness_coeffici entbdticm(- 0.594136) | length(0.491)evergreen_ne edleleaf_tree(0.10 6) lengthgeol_porosit | shape_factortksatu_1 6(-0.372096) compactness_coefficie |
|--|--|---|--|---|--|
| <u>i</u> | 95) | 51 <u>203</u>) | mt(0.566199) | <u>y(-0.451183)</u> | nt(-0.421)va(0.172) |
| elev(m)Va | prs_mean(pa(0.889271) | e_mingeol_porosity(-0.753219) | lon(-yb(0. 752 21) | e_max(0.752)deciduous_n eedleleaf_tree(0.1 86) | theta_s_14(pi(0.7172) |
| s lope(m/km)wb | <u>n_min(</u> <u>water_bodies(</u> 0.552 39) | permanent_wetland (0.551264) | n_max(bldfie_sl4(0.5514 8) | phihox_sl7(bldfie_sl5(0.49114 7) | oredre_sll_urban_an d_built- up_land(0.49138) |
| n_minpb | lat(1.mt(0.176) | frac_snow_dailypa(0.7 | <u>gst_meantheta_s_l</u> <u>5(-0.693128)</u> | pre_mean(- area(0.651 <u>127</u>) | tem_mean(- length(0.648123) |
| n_maxvb | lat(1.va(0.21) | frac_snow_daily(geol_ porosity(-0.701171) | gst_mean(vi(0.692165) | pre_mean(cecsol_sl7(0.6s16 1) | tem_mean(cecsol_sl6(0.647 <u>15</u> <u>7</u>) |
| e_minnd | lon(1.barren(0.154) | elev(aridity(0.753146) | evp <u>pre</u> _mean(- 0.731 <u>144</u>) | prs_mean(lai_dif(- 0.707141) | ndvi_meanSnow_an d_ice(0.691141) |
| e_maxpy | lon(1.phihox_sl1(- 0.237) | elevphihox_sl2(- 0.752233) | evp_meanphihox_s 13(-0.729233) | prs_mean(phihox_sl 4(-0.70723) | ndvi_meanwoody_s avanna(0.69227) |
| pop(people) <u>ev</u> | ************************************** | urban and built up land(0.399)orcdrc_sl5 (-0.035) | tem_mean(orcdrc_s 14(-0.318035) | p_seasonality(cecsol _sl3(-0.317034) | frae_snow_dailyorcdr c_s17(-0.304034) |
| tksatu 11pop_dnsty(people/km²) | urban and built up land(0.811)grav(- 0.346) | p_seasonality(som(- 0.426 <u>344</u>) | tem_meanbldfie_sl 3(0.412298) | gst_meanbldfie_s11(0.395295) | frae_snow_daily(bldfie_sl2(0.39291) |
| tksatu 121 0n | e_max(1-som(- 0.365) | e_min(1.bldfie_sl3(0. 326) | bldfie_sl1(0.7523 26) | evp_mean(- bldfie_sl2(0.73323) | prs_mean (grav(- 0.707308) |
| lattksatu_13 | n_min(1.som(-0.344) | n_max(1.bldfie_sl2(0 .328) | frac_snow_dailybldf ie_sl1(0.702325) | <u>gst_mean(</u> <u>bldfie_sl3(</u> 0.69332 <u>4</u>) | <u>pre_mean(</u> <u>bldfie_sl4(</u> 0.65130 <u>8</u>) |

| tksatu_ <u>11<u>14</u></u> | snow and icebldfic_s12(0.5033 98) | siltsom(-0.465397) | som(bldfie_sl1(0.3663 88) | sandbldfie_sl3(0.36 2384) | \frac{\log_k_s_15\text{bldfie_sl}}{4(0.327\frac{358}{})} |
|----------------------------------|--|---|--|--|--|
| tksatu_ <u>12</u> <u>15</u> | ************************************** | silt(- bldfie_sl2(0.49376) | sand(som(- 0.406369) | som(- bldfie_sl4(0.36536 4) | log_k_s_I\$bldfie_sl 1(0.364358) |
| tksatu_ 13 <u>l6</u> | ************************************** | som silt (-0.489 <u>362</u>) | sandbdticm(0.4093 6) | ndvi_mean(bldfie_sl2(0.36834 3) | elay(bldfie_sl7(0.33433 8) |
| tksatu_14 <u>log_k</u> s_11 | snow and icesand(0.53371) | siltclay(-0.49 <u>59</u>) | sand(savanna(- 0.465 <u>441</u>) | ndvi_meansilt(- 0.455 <u>436</u>) | log_k_s_15(rhu_mea n(-0.414423) |
| tksatu_15log_k_ s_12 | snow and icesand (0.568709) | sitclay(-0.402 <u>578</u>) | ndvi_meansavanna (-0.375452) | sandphihox_sl7(0.3 48438) | lai_difsilt(- 0.326433) |
| tksatu_l6log_k_ | snow and | bdtiem(clay(- | log_k_s_l6(savanna | suphihox_s17(0.38 | low_prec_freqphihox |
| <u>s_13</u> | ice sand(0.449 <u>682</u>) | 0.4 03 592) | <u>(-</u> 0. 384<u>448</u>) | <u>442</u>) | <u>sl6</u> (0. 36 435) |
| log_k_s_ <u>11114</u> | sand(0. <u>858612</u>) | clay(-0. 733<u>603</u>) | pre_mean_Savanna(- 0.55349) | phihox_sl7(pre_mea n(-0.551489) | phihox_s16s17(0.54 6485) |
| log_k_s_ 12 <u>15</u> | sand(clay(-0.86561) | elay(_sand(0.729 <u>555</u>) | phihox_sl7(0. 575 <u>506</u>) | phihox_sl6(savanna(_0.569501) | pre_mean(phihox_sl6(0.5685 01) |
| log_k_s_ 13 <u>16</u> | sand(clay(-0.859563) | elaypre_mean(- 0.728 <u>555</u>) | pre_mean(aridity(0.571548) | phihox_sl7(0. 571 <u>5</u> <u>34</u>) | phihox_sl6(0. 565 <u>5</u> <u>32</u>) |
| log_ktheta_s_14 11 | sand(grav(-0.82582) | clay <u>←(</u> 0. 752 <u>325</u>) | pre_meansand(- 0.647 <u>315</u>) | phihox_sl7(elev(- 0.636314) | phihox_sl6pdep(0.63 311) |
| log_ktheta_s_15 12 | sand(grav(- 0.773 <u>585</u>) | elay(-pdep(0.714377) | phihox_sl7(elev(- 0.654 <u>366</u>) | clayphihox_sl6(0. | phihox_sl5(sand(- 0.646326) |
| log_ktheta_s_l6 13 | grav(- sand(0.688 <u>522</u>) | elay(pdep(0.68742) | phihox_sl7(elev(- 0.665414) | phihox_sl6prs_mean (0.662365) | pre_mean(clay(0.662359) |
| theta_s_111 <u>14</u> | grav(-0. 705 <u>515</u>) | elev(-pdep(0.422463) | rhu_mean(clev(- 0.407 <u>412</u>) | elayprs_mean(0.40 1349) | pdeplon(0.4 <u>328</u>) |
| theta_s_ 12 <u>15</u> | grav(-0. 713433) | elev(-0. 505401) | pdep(0.475 <u>376</u>) | e_min(sand(- 0.442349) | lonrhu_mean(0.441 331) |
| - | | | | | |

| - | grav(- | | | | |
|------------------------------------|---|--|--|--|---|
| theta_s_ 13 <u>16</u> | 0.662) evergreen_bro | elevgrav(-0.638357) | prs_mean(elev(- | pdep(sand(- | $\frac{e_{min}tem_mean}{0}$. |
| | adleaf tree(0.372) | 0.030 <u>337</u>) | 0. 55 4 <u>344</u>) | 0. 52 343) | 516 337) |
| theta s 14orcdr | elevbldfie sl4(- | gravbldfie_sl5(- | prs_mean(bldfie_sl | pdep(bldfie sl3(- | e min(bldfie sl7(- |
| <u>c sl7</u> | 0.7 <u>581</u>) | 0. 663 <u>572</u>) | 6(-0.594548) | 0. 571 535) | 0. 51 523) |
| | | | prs mean(bldfie sl | pdep(bldfie sl4(- | rhu mean(bldfie sl |
| theta_s_15orcdr | elevbldfie_sl3(- | gravbldfie_sl2(- | | | |
| <u>c_sl3</u> | 0.656 <u>738</u>) | 0.584 <u>728</u>) | <u>1(-</u> 0. 536 <u>701</u>) | 0. 501 <u>691</u>) | <u>5(-</u> 0.467 <u>621</u>) |
| theta_s_16orcdr | elevbldfie_sl3(- | prs_mean(bldfie_sl2(| gravbldfie_sl4(- | bldfie_sl1high_pr | rhu_mean(bldfie_sl |
| <u>c_sl4</u> | 0. 637<u>702</u>) | <u>-</u> 0. 525 <u>682</u>) | 0. 513 <u>676</u>) | ee_dur (-0. 503 <u>657</u>) | <u>5(-</u> 0.4 75 <u>614</u>) |
| orcdrc_ sl7 s <u>l5</u> | cccsol_sl2(bldfie_sl4 | bldfie_ <u>sl2sl3</u> (- | bldfie_ <u>s44s12</u> (- | bldfie_ <u>sHsl5</u> (- | cccsol_sl3(bldfie_sl |
| 010d10_31/ <u>515</u> | <u>(-</u> 0. 758 <u>641</u>) | 0. 745<u>636</u>) | 0. 744<u>611</u>) | 0. 737 <u>6</u>) | <u>1(-</u> 0. 735 <u>592</u>) |
| | bldfie_ <u>sl2sl4</u> (- | bldfie_ <u>sl4sl5</u> (- | bldfie_s <u>l3sl6</u> (- | bldfie_ <u>sl5</u> s <u>l3</u> (- | bldfie_s11s17(- |
| orcdrc_ sl3 sl6 | 0. 876<u>584</u>) | 0. 875 <u>567</u>) | 0. 874<u>556</u>) | 0. 849<u>552</u>) | 0. 848<u>534</u>) |
| 1 14-12 | bldfie_ <u>sl4</u> s <u>l2</u> (- | bldfie_ sl2 sl1(- | bldfie_sl3(- | bldfie_ <u>sl5s14</u> (- | bldfiececsol_sl1(- |
| orcdrc_ s 14 <u>s12</u> | 0. 823 <u>787</u>) | 0. 809 769) | 0. 803<u>749</u>) | 0. 803 <u>68</u>) | <u>(</u> 0. 787 <u>629</u>) |
| anadna 15ali | bldfie_sl4phihox_sl2 | bldfie_sl2phihox_sl3 | bldfie_sl5phihox_s | bldfie_sl1phihox_sl | bldfie_sl3phihox_sl |
| orcdrc_ sl5 <u>s11</u> | (-0. 759 <u>599</u>) | (-0. 75 4 <u>594</u>) | <u>14</u> (-0. 745 <u>591</u>) | <u>5</u> (-0. 745 <u>586</u>) | <u>6</u> (-0. 731 <u>585</u>) |
| 1 tolding | woody_savanna(- | 1116 14 | 1116 127 | bldfie_sl1(- | |
| oredre_sl6phiho | eccsol_sl2(0.733 <u>628</u> | bldfie_sl4pre_mean(- | bldfie_sl2(aridity(0.728592) | low_prec_freq(0.7 | orcdrc_sl1bldfie_s |
| <u>x_sl7</u> |) | 0. 733<u>598</u>) | <u>andity(</u> 0. 728 392) | 25 588) | 15 (-0. 721<u>583</u>) |
| phihox sl6ore | bldfie_sl2woody_sav | bldfie_sl1pre_mean(- | bldfie_sl3(- | cecsol_sl1(lai_max(| bldfie_sl4orcdrc_sl1 |
| drc_sl2 | <u>anna</u> (-0.917 <u>627</u>) | 0. 908<u>594</u>) | aridity(0.86159) | <u>-</u> 0. 85 4 <u>587</u>) | (-0. 854<u>585</u>) |
| phihox sl5ore | phihox_sl2woody_sa | phihox_sl1lai_max(- | phihox_sl3pre_mea | phihox_sl4(- | phihox_sl5orcdrc_sl |
| dre_sl1 | <u>vanna</u> (-0.826626) | 0.824 <u>593</u>) | <u>n</u> (-0. <u>822</u> 592) | aridity(0.819589) | <u>1</u> (-0. <u>813586</u>) |
| phihox_s 17 s <u>14</u> | low_prec_freq(woody | 10: | 1 | 1 111-: 1:27 | 1: 1:0 |
| | <u>savanna(-</u> | pre_meanlai_max(- | lai_maxorcdrc_sl1 | oredre_sl1 <u>lai_dif(</u> - | lai_difpre_mean(- |
| | 0. 825 <u>628</u>) | 0. 819<u>599</u>) | (-0. 806<u>591</u>) | 0. 804<u>578</u>) | 0. 799 <u>576</u>) |
| phihox_sl6s13 | low_prec_freq(woody | | | 1 111-: 1:27 | 1: 1:0 |
| | _savanna(- | lai_max(-0. 814<u>595</u>) | pre_meanorcdrc_sl | oredre_sl1_laidif(- | lai_difpre_mean(- |
| | 0. 818<u>627</u>) | | <u>1</u> (-0. 81 <u>594</u>) | 0. 807<u>576</u>) | 0. 807 <u>568</u>) |
| phihox_ sl5 s <u>l2</u> | lai max woody sava | low prec freq(lai ma | orcdrc_sl1(- | | pre_mean(- |
| | | | | lai_dif(-0. 807 <u>583</u>) | low_prec_freq(0.8 |
| | <u>nna</u> (-0. 814<u>627</u>) | <u>x(-</u> 0. 81 4 <u>602</u>) | 0. 813 <u>599</u>) | | 01 <u>569</u>) |
| | | | | | |

| phihox_s44s11 | oredre_sll_woody_sa vanna(-0.819601) | lai_max(-0. <u>815<u>586</u>)</u> | lai_diforcdrc_sl1(- 0.809584) | low_prec_freq(lai_di f(-0.804565) | pre_mean(bldfie_s12(0.78155) |
|------------------------------------|---|---------------------------------------|---|---|--|
| phihox_sl3bldfi e_sl7 | orcdrc_s <u>Hsl5</u> (- 0.822 <u>547</u>) | lai_maxorcdrc_sl4(-0.813546) | lai_diforcdrc_sl3(-0.806543) | orcdrc_sl6(- low_pree_freq(0.7 99534) | pre_meanorcdrc_sl7 (-0.772523) |
| phihox_sl2bldfi | orcdrc_s11s15(- | lai_max orcdrc_sl6(- | lai_diforcdrc_sl4(- | low_prec_freq(orcdr | pre_meanorcdrc_sl3 |
| <u>e_sl6</u> | 0. 826<u>559</u>) | 0. 813<u>556</u>) | 0. 807<u>553</u>) | <u>c_s17(-</u> 0. 798 <u>548</u>) | (-0. 767<u>547</u>) |
| bldfie sl5phiho | orcdrc_ s11 <u>s13</u> (- | lai_max orcdrc_sl4(- | lai_diforcdrc_sl5(- | low_prec_freq(orcdr | pre_meanorcdrc_sl7 |
| x_s11 | 0. 82 4 <u>621</u>) | 0. 80 4 <u>614</u>) | 0. 798 <u>6</u>) | <u>c_sl2(-</u> 0. 78 <u>597</u>) | (-0. 741<u>572</u>) |
| bldfie s17s14 | orcdrc_sl3(- | orcdrc_ sl 4 <u>sl2</u> (- | orcdrc_ sl5 <u>sl4</u> (- | orcdrc_ sl2 sl5(- | orcdrc_sl6(- |
| 01d11e_ s1/ <u>514</u> | 0. 775 <u>691</u>) | 0. 747<u>68</u>) | 0. 698 676) | 0. 698<u>641</u>) | 0. 671 <u>584</u>) |
| bldfie s16s11 | orcdrc_ <u>sl3s12</u> (- | orcdrc_ <u>s14</u> <u>s13</u> (- | oredre_sl5cecsol_sl | orcdrc_ <u>sl2s14</u> (- | oredre_sl6som(- |
| bidite_ sio sii | 0. 776 <u>769</u>) | 0. 748<u>701</u>) | <u>1</u> (-0. 701<u>686</u>) | 0. 69 4 <u>657</u>) | 0. 677<u>606</u>) |
| bldfie s15s13 | orcdrc_ sl3 <u>sl2</u> (- | orcdrc_ <u>s12s13</u> (- | orcdrc_sl4(- | orcdrc_sl5(- | oredre_sl7som(- |
| olune_ si3 si3 | 0. 849<u>749</u>) | 0.81738) | 0. 803 <u>702</u>) | 0. 745 <u>636</u>) | 0. 728<u>633</u>) |
| 1110 12 | orcdrc_ <u>sl3s12</u> (- | orcdrc_ <u>s12</u> s <u>13</u> (- | orcdrc_sl4(- | cecsol_sl1(- | oredre_sl5som(- |
| bldfie_ sl 4 <u>sl2</u> | 0. 875 <u>787</u>) | 0. 854<u>728</u>) | 0. 823 <u>682</u>) | 0. 763 <u>671</u>) | 0. 759<u>651</u>) |
| bldfiececsol_sl | oredre_sl2bldfie_sl1(| eecsol_s11bldfie_s12(| orcdrc_ sl3(- | eecsol_sl2bldfie_sl3 | orcdrc_ sl4(|
| 1 | -0. 908<u>686</u>) | -0. 891<u>671</u>) | <u>sl2(</u> 0. 848 629) | (-0. 828 <u>598</u>) | <u>sl3(</u> 0. 787 <u>579</u>) |
| cecsol_sl2 _{bldfi} | oredre_sl3bldfie_sl1(-0.874 <u>579</u>) | oredrebldfie_sl2(-0.861566) | orcdrc_s14(- s12(0.803553) | eecsol_sl1(oredre_sl3(0.7955 23) | sombldfie_sl3(- 0.787515) |
| bldfie_sl2cecsol | oredre_sl2bldfie_sl1(-0.917445) | oredre_sl3bldfie_sl2(-0.876429) | eecsol_sl1(orcdrc_sl2(0.874 12) | oredre_sl4(- sl3(0.809393) | sompet_mean(- 0.808392) |
| cecsol_s11s14 | bldfie_sl1(- | bldfie_sl2(- | orcdrc_sl2(0. 854 <u>4</u> | bldfieorcdrc_sl3(- | orcdrc_ sl3 <u>sl5</u> (0. 781 |
| | 0. 891<u>472</u>) | 0. 87<u>459</u>) | <u>47</u>) | <u>(</u> 0. 795 <u>43</u>) | <u>424</u>) |
| cecsol_s12s13 | bldfie_sl1(- | oredrebldfie_sl2((- | bldfieorcdrc_sl2(- | orcdrc_ <u>sl7s13</u> (0. 758 | orcdrc_s13s14(0.746 |
| | 0. 828<u>532</u>) | 0. 822 <u>52</u>) | (0. 798 <u>508</u>) | <u>49</u>) | <u>478</u>) |
| cecsol_ sl5 sl7 | bldfie_sl1(- | oredrebldfie_sl2((- | orcdrc_ <u>s17s12</u> (0.64 | bldfie_sl2pet_mean(| orcdrc_s16s13(0.636 |
| | 0. <u>681<u>413</u>)</u> | 0. 66 4 <u>396</u>) | 9 <u>38</u>) | -0. 645<u>374</u>) | <u>362</u>) |
| cecsol_s14s16 | bldfie_sl1(- | oredrebldfie_sl2((- | orcdrc_ s17 <u>s12</u> (0. 69 | bldfie_sl2pet_mean(| orcdrc_s16s13(0.679 |
| | 0. 72 409) | 0. 717<u>393</u>) | 3 <u>378</u>) | -0. 692 <u>373</u>) | <u>36</u>) |
| | | | | | |

| eecsol_sl3bdtic | bldfie_sl1(su(0.784 <u>52</u>) | oredre_sl2(woody_sa vanna(-0.776412) | bldfie_sl2(low_prec_freq(0. 76382) | oredrephihox_sl7(0 .735378) | orest(-0.733374) |
|----------------------------|---|--|--|--|--|
| eeesol_sl7pdep | bldfie_sl1(theta_s_14(0.661463) | oredre_s17(elev(- 0.654 <u>436</u>) | o redre_sl2(grav(- 0.64 <u>2424)</u> | oredre_sl6theta_s_1 3(0.6442) | oredre_s15 <u>lon</u> (0.619 <u>4</u>) |
| cecsol_sl6 por | bldfie_sl1(som(0.648363) | oredre_sl2(bldfie_sl1 (-0.637335) | oredre_sl7(phihox_sl1(-0.632329) | oredre_sl6(phihox_s 13(-0.62328) | bldfiephihox_sl2(- 0.61328) |
| <u>clay</u> bdtiem | su(sand(-0.59967) | low_prec_freq(log_k s_14(-0.463603) | log_k_s_ 16(13(- 0.439 <u>592</u>) | phihox_sl2(log_k_s _11(-0.43759) | phihox_s17(log_k_s _12(-0.436578) |
| sand _{pdep} | elev(log k s 11(0.66271) | thetalog_k_s_1412(0. | e_minlog_k_s_13(0.566682) | lon(clay(-0.56567) | e_maxlog_k_s_14(0 .564612) |
| <u>silt</u> por | silt(sand(-0.573) | log k s 11(- elay(0.366436) | tksatulog_k_s_12(- 0.317433) | som(log_k_s_13(- 0.3144) | tksatu_11log k s 14 (-0.309316) |
| clay grav | pre_mean(theta_s_12(-0.763585) | log_ktheta_s_14 <u>11</u> (- 0.752 <u>582</u>) | log_ktheta_s_1113(-0.733522) | log_ktheta_s_1214(- 0.729515) | log_ktheta_s_1315(- 0.728433) |
| sandSOM | log_k_s_12(bldfie_sl2 (-0.86651) | log_k_s_13(bldfie_sl3 (-0.859633) | log_k_s_H(bldfie_s 11(-0.858606) | log_k_s_14orcdrc_sl 2(0.82599) | \frac{\log_k_s_15\text{orcdrc_sl}}{3(0.773\text{576})} |
| sithigh prec f | por(root_depth_50(- 0.573196) | sand(grassland(0.558175) | bg_k_s_l3root_de pth_99(- 0.557171) | log_k_s_12(som(0.547136) | log_k_stksatu_11(- 0.545133) |
| high prec du | theta_s_ <u>1216</u> (- 0. 713 <u>277</u>) | theta_s_ <u>++15</u> (- 0. 70 \$234) | theta_s_14(- p_seasonality(0.6 63233) | theta_s_13(elev(0.662211) | theta_s_ <u>1514</u> (- 0. <u>584201</u>) |
| low prec fre | bldfie_sl2pre_mean(- 0.808766) | bldfie_sl3(aridity(0.787745) | <u>ssd_mean(0.7596</u> <u>52</u>) | bldfie_sl4 <u>rhu_mean</u> (-0.747 <u>627</u>) | oredre_sl2phihox_sl 7(0.74588) |

Appendix C: Data sources and data processing

The program to generate the data set is mainly written in Python. The rasterio¹³ library is used to extract from the raster for the given basin boundary, reproject and merge rasters: The shapely 14 library is used to calculate the geometry: The pyproj 15 library is used for coordinate system conversions: The richdem 16 library is used to calculate slope: The netCDF4 17 and xarray 18 library is used to read the netCDF files; The pyshp¹⁹ library is used to handle shapefiles; The gdal²⁰ command-line programs 620 are used for data format conversions: The Python multiprocessing 21 library is used for multi-threaded data processing such as the calculation of meteorological time series: The interpolation program is written based on SciPv and NumPv. In addition, the calculation of the catchment boundary uses ArcPy ²². However, ArcPy is not open sourced. The dataset can SURF CLI CHN MUL DAY be downloaded 625 https://data.cma.cn/data/cdcdetail/dataCode/SURF CLI CHN MUL DAY.html. It is freely available to global researchers but registration is required. The GDBD dataset can be downloaded at https://www.cger.nies.go.ip/db/gdbd/gdbd index e.html. ASTER GDEM dataset can be downloaded at: https://asterweb.jpl.nasa.gov/gdem.asp. GLHYMPS dataset can be downloaded at: https://dataverse.scholarsportal.info/dataset.xhtml?persistentId=doi:10.5683/SP2/DLGXYO; MODIS MCD12O1 can be obtained from: https://lpdaac.usgs.gov/products/mcd12q1v006/; MODIS MCD15A3 can be obtained from: https://lpdaac.usgs.gov/products/mcd15a3hv006/; Soil hydraulic and thermal properties can be downloaded after registration: 630 http://globalchange.bnu.edu.cn/research/soil5.jsp; Soil properties data can be downloaded after registration: http://globalchange.bnu.edu.cn/research/soil2; SoilGrids250m data download links: https://files.isric.org/soilgrids/former/2017-03-10/data/ with list of descriptions: https://github.com/ISRICWorldSoil/SoilGrids250m/blob/master/grids/models/META_GEOTIFF_1B.csv.

635 Appendix D: Basin boundaries

This section briefly introduces how the basin boundaries are derived. The basin boundaries data used in this research are obtained from the GBDB (Masutomi, Inui et al. 2009) dataset. The GDBD dataset first distinguishing sinks caused by DEM

¹³ https://rasterio.readthedocs.io/en/latest/

¹⁴ https://shapely.readthedocs.io/en/stable/manual.html

¹⁵ https://pyproj4.github.io/pyproj/stable/

¹⁶ https://richdem.readthedocs.io/en/latest/

¹⁷ https://unidata.github.io/netcdf4-python/

¹⁸ http://xarray.pydata.org/en/stable/

¹⁹ https://pypi.org/project/pyshp/

²⁰ https://gdal.org/api/python.html

²¹ https://docs.python.org/3/library/multiprocessing.html

²² https://pro.arcgis.com/zh-cn/pro-app/latest/arcpy/get-started/what-is-arcpy-.htm

errors, then the stream burning (Maidment 1996), and ridge fencing methods are used to modify the seeded DEM, then basin boundaries are produced with standardized procedures (Jenson, Domingue et al. 1988, Maidment and Morehouse 2002). Then the gauging station data from the GRDC (Center 2005) dataset is used to calibrate the derived basin boundaries. The derived basin areas were compared with the observed basin areas, and they showed a high degree of consistency with the observed basin data.

Appendix E: Guidelines for generating basin attributes for any basin

The published code²³ supports the automation of the calculation of the attributes for any given river basin and the generation of statistics files. In general, the user only needs to prepare the source data and ensure that the code environment is installed correctly, and then the user can run the code to calculate all attributes for the given river basin. The following describes the steps to generate data for any given watershed.

Prepare source data

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In this step, the user needs to download the source data and place it in the corresponding location (Table D1). The code supports the calculation of meteorological time series based on the SURF CLI CHN MUL DAY data set. If the basin the user need to calculate is not in China, then the user needs to format the collected meteorological time series into the same format as the time series generated by the code. A sample file is available in the GitHub library.

655 Table D1: Instructions for preparing data sources

| Data | <u>Download link</u> | Example | Note |
|---------------|-------------------------------------|--|------|
| source | | | |
| <u>ASTER</u> | https://search.earthdata.nasa.gov/ | ./data/dems/ *.tif | |
| <u>GDEM</u> | search/ | | |
| | https://www.jspacesystems.or.jp/ | | |
| | ersdac/GDEM/E/ | | |
| <u>GLHYMP</u> | https://dataverse.scholarsportal.in | <pre>./data/processed_permeability</pre> | |
| <u>S</u> | fo/dataset.xhtml?persistentId=doi | <u>.tif</u> | |
| | :10.5683/SP2/DLGXYO (using | <pre>./data/processed_porosity.tif</pre> | |
| | source data requires merging | | |
| | multiple small pieces to a single | | |
| | TIFF) | | |

²³ https://github.com/haozhen315/CCAM-China-Catchment-Attributes-and-Meteorology-dataset

| | https://ldrv.ms/u/s!AqzR0fLyn9 | | |
|-------------|---------------------------------------|--------------------------------------|-------------------------------------|
| | KKspF6HAAuXU9Twkkz1Q?e= | | |
| | QCPFAm (our processed file) | | |
| | https://ldrv.ms/u/s!AqzR0fLyn9 | | |
| | $\underline{KKspF70EPmDubS5V2qTQ?e=}$ | | |
| | Rbybwa (our processed file) | | |
| <u>GLiM</u> | https://csdms.colorado.edu/wiki/ | ./data/processed_glim.py | |
| | <u>Data:GLiM</u> | | |
| | https://1drv.ms/u/s!AqzR0fLyn9 | | |
| | KKspF5Vktb- | | |
| | zlmd_Ctxg?e=G6fOuh (our | | |
| | processed file) | | |
| MCD12Q1 | https://lpdaac.usgs.gov/products/ | <pre>./data/processed_igbp.tif</pre> | |
| | mcd12q1v006/ | | |
| | https://1drv.ms/u/s!AqzR0fLyn9 | | |
| | KKspF4xxbe0xM7qJNzkA?e=vy | | |
| | FcFj (our processed file) | | |
| MCD15A3 | https://lpdaac.usgs.gov/products/ | ./data/MCD15A3/ | |
| | mcd15a3hv006/ | MCD15A3H.A2002185.h22v | |
| | | 04.006.2015149102803.hdf | |
| MOD13Q1 | https://lpdaac.usgs.gov/products/ | ./data/MOD13Q1/MOD13Q1 | |
| | mod13q1v006/ | .A2002186.h22v04.006.2015 | |
| | | <u>149102803.hdf</u> | |
| <u>Soil</u> | http://globalchange.bnu.edu.cn/re | ./data/soil_souce_data/binary/ | |
| | search/soil5.jsp | <u>log k s 11</u> | |
| <u>Soil</u> | https://files.isric.org/soilgrids/for | ./data/soil_souce_data/tif/BD | Description: |
| | mer/2017-03-10/data/ | TICM_M_250m_ll.tif | https://github.com/ISRICWorldSoil/S |
| | | | oilGrids250m/blob/master/grids/mode |
| | | | ls/META_GEOTIFF_1B.csv |
| <u>Soil</u> | http://globalchange.bnu.edu.cn/re | ./data/soil_souce_data/tif/SA. | |
| | search/soil2 | <u>nc</u> | |

| SURF_CLI | https://data.cma.cn/data/cdcdetail | ./data/SURF_CLI_CHN_MU | If basin boundary is outside China, |
|-------------|------------------------------------|--------------------------------|---|
| _CHN_M | /dataCode/SURF_CLI_CHN_M | L_DAY/Data/EVP/SURF_C | format and place the collected time |
| UL_DAY | UL_DAY.html | LI CHN MUL DAY-EVP- | series data in |
| | | <u>13240-195101.TXT</u> | ./output/catchment_meteorological |
| Root depth | https://github.com/haozhen315/C | ./data/root_depth_calculated.t | Calculated root depth of each land type |
| | CAM-China-Catchment- | <u>xt</u> | according to (Zeng 2001). |
| | Attributes-and-Meteorology- | | |
| | dataset/blob/main/data/root_dept | | |
| | h_calculated.txt | | |
| <u>GLiM</u> | https://github.com/haozhen315/C | ./data/glim_cate_number_ma | These files are used for name |
| name | CAM-China-Catchment- | pping.csv | conversions in the program. |
| mapping | Attributes-and-Meteorology- | ./data/glim_name_short_long. | |
| | dataset/blob/main/data/glim_nam | <u>txt</u> | |
| | e_short_long.txt | | |
| | https://github.com/haozhen315/C | | |
| | CAM-China-Catchment- | | |
| | Attributes-and-Meteorology- | | |
| | dataset/blob/main/data/glim_cate | | |
| | number_mapping.csv | | |
| <u>GDBD</u> | https://www.cger.nies.go.jp/db/g | _/data/river_network/as_strea | River network shapefiles are used to |
| | dbd/gdbd_index_e.html | ms_wgs.shp | determine river basin shape factors. |
| | | | The source data need to be reprojected |
| | | | to EPSG:4326 (using ArcMap or |
| | | | QGIS) to successfully run the code. |
| | | | Note that files in different regions have |
| | | | different names. |

Run the code

When all the data is ready, the user can run the code calculate_all_attributes.py to calculate all attributes or run separate scripts (e.g., soil.py) to calculate indicators for specific categories. The result will appear in the output folder.

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