1	
1	Landsat and Sentinel-derived glacial lake dataset in the
2	China-Pakistan Economic Corridor from 1990 to 2020
3	
4	Muchu Lesi ¹ , Yong Nie ^{1, *} , Dan H. Shugar ² , Jida Wang ³ , Qian Deng ^{1, 4} , Huayong Chen ¹
5	¹ Institute of Mountain Hazards and Environment, Chinese Academy of Sciences, Chengdu,
6	China.
7	² Water, Sediment, Hazards, and Earth-surface Dynamics (waterSHED) Lab, Department of
8	Geoscience, University of Calgary, Alberta, T2N 1N4, Canada
9	³ Department of Geography and Geospatial Sciences, Kansas State University, Manhattan,
10	Kansas 66506, USA
11	⁴ University of Chinese Academy of Sciences, Beijing 100190, China
12	
13	
14	
15	*Corresponding author, nieyong@imde.ac.cn
16 17	

18 Abstract. The China-Pakistan Economic Corridor (CPEC) is one of thea number of flagship 19 projects of the One Belt One Road Initiative, which faces threats from mountain disasters in 20 the high altitude region, such as glacial lake outburst floods (GLOFs). An up-to-date high-quality glacial lake dataset with critical-parameters likesuch as lake type, acquisition 21 date and uncertainty area (e.g. lake types), which is fundamental to flood risk assessments and 22 23 predicting glacier-lake evolutions and cryosphere-hydrological interactionschanges, is still 24 largely absent for the entire CPEC. This study describes a glacial lake dataset in 2020 for CPEC, based on an object-oriented mapping method associated with rigorous visual 25 inspection workflows. This dataset includes (1) a glacial lake inventory for the year 2020 at 26 10 m resolution produced, which was produced from Sentinel optical spectral images, and (2) 27 multi-temporal inventories for 1990, 2000, and 2020 produced from as well as 30 m 28 resolution Landsat imagesglacial lake inventories in 1990, 2000 and 2020 from Landsat-29 30 observation at 10-30 m resolution, which was produced from both Landsat and Sentineloptical images as well as glacial lake inventories in 1990 and 2000 from Landsat observation. 31 32 using an advanced object-oriented mapping method associated with rigorous visualinspection workflows. The results show that Landsat derived 2234 glacial lakes in 2020, 33 covering a total area of $86.31 \pm 14.98 \text{ km}^2$ with a minimum mapping unit of 5 pixels (4500 m²), 34 whereas Sentinel derived 7560 glacial lakes in 2020 with a total area of 103.70 ± 8.45 km² 35 with a minimum mapping unit of 5 pixels (500 m^2). The discrepancy implies that there is a 36 significant quantity of small glacier lakes not recognized in existing glacial lake inventories 37 and a more thorough inclusion of them require future efforts using higher resolution data. The 38 total number and area of glacial lakes from consistent 30 m resolution Landsat images remain 39 40 relatively stable despite a slight increase from 1990 to 2020. A range of critical attributes 41 have been generated in the dataset, including lake types of two classification systems and 42 mapping uncertainty estimated by an improved Hanshaw's equation. This comprehensive 43 glacial lake dataset has potentials to be widely applied in studies on glacial lake-related hazards-and, glacier-lake interactions and cryospheric hydrology, and is freely available at 44 45 https://doi.org/10.12380/Glaci.msdc.000001 (Lesi et al., 2022) (Lesi et al., 2022).-

46 **1 Introduction**

47 Glaciers in High-mountain Asia (HMA) play a crucial role in regulating climate, supporting 48 ecosystems, modulating the release of freshwater into rivers, and sustaining municipal water 49 supplies (Wang et al., 2019; Viviroli et al., 2020) (Wang et al., 2019; Viviroli et al., 2020), 50 agricultural irrigation, and hydropower generation (Pritchard, 2019; Nie et al., 2021)-(Pritchard, 2019; Nie et al., 2021). Most HMA glaciers are losing mass in the context of 51 52 climate change (Brun et al., 2017; Maurer et al., 2019; Shean et al., 2020; Bhattacharya et al., 2021) (Brun et al., 2017; Shean et al., 2020; Bhattacharya et al., 2021; Maurer et al., 2019), 53 54 therefore, unsustainable glacier melt is reducing the hydrological role of glaciers and 55 impacting downstream ecosystem services, agriculture, hydropower and other socioeconomic values (Carrivick and Tweed, 2016; Nie et al., 2021) (Nie et al., 2021). The present and 56 57 future glacier changes also alter the frequency and intensity of glacier-related hazards, such as glacier lake outburst floods (GLOFs) (Nie et al., 2018; Rounce et al., 2020; Zheng et al., 58

2021) (Nie et al., 2018; Zheng et al., 2021; Rounce et al., 2020), and rock and ice avalanches_
(Shugar et al., 2021) (Shugar et al., 2021). Global glacial lake numbers and total area both_
increased between 1990 and 2018 in response to glacier retreat and climate change (Shugar et al., 2020), which inevitably affected the risk of GLOFs., altering the risk of GLOFs. The
increasing frequency of GLOFs has been observed in the Karakoram and Himalaya (Nie et al., 2020)

64 <u>2021</u> (Nie et al., 2021), and the increasing risk of GLOFs (Zheng et al., 2021) is threatening

existing and planned infrastructures in the mountain ranges, such as hydropower plants,railways, and highways.

67 A large number of major infrastructure construction projects for the One Belt One Road 68 Initiative (BRI) play a fundamental role in strengthening the interconnection of infrastructure 69 between countries and promoting international trade and investment (Battamo et al., 2021; Li 70 et al., 2021) (Battamo et al., 2021; Li et al., 2021). Taking the Karakoram Highway for 71 example, it is a unique land route to link China and Pakistan. The China-Pakistan Economic 72 Corridor (CPEC) is one of the BRI flagship projects, originating from Kashgar of the 73 Xinjiang Uygur Autonomous region, China and extending to Gwadar Port, Pakistan (Ullah et al., 2019; Yao et al., 2020)-(Ullah et al., 2019; Yao et al., 2020). The northern section of the 74 75 CPEC passes through Pamir, Karakoram, Hindu Kush and Himalaya mountains where glacier-related hazards such as GLOFs are frequent and severe (Hewitt, 2014; Bhambri et al., 76 2019) (Hewitt, 2014; Bhambri et al., 2019), threatening the existing, under-construction and 77 78 planned infrastructure projects. Understanding the risk posed by GLOFs is a critical step to 79 disaster prevention for infrastructures across the CPEC (Figure 1Figure 1).

80 Glacial lake inventories with a range of attributes benefit risk assessment and disaster 81 reduction related to GLOFs, and contribute to predicting glacier-lake evolution and 82 cryosphere-hydrosphere interactions under climate change (Nie et al., 2017; Brun et al., 2019; Maurer et al., 2019; Carrivick et al., 2020; Liu et al., 2020) (Nie et al., 2017; Brun et al., 2019; 83 84 Liu et al., 2020; Maurer et al., 2019). Remote sensing is the most viable way to map glacial lakes and detect their spatio-temporal changes in the high-elevation zones where in situ 85 accessibility is extremely low (Huggel et al., 2002; Quincey et al., 2007) (Huggel et al., 2002; 86 87 Quincey et al., 2007). Studies in glacial lake inventories using satellite observations have been heavily conducted at regional scales recently, such as in the Tibetan Plateau (Zhang et 88 al., 2015) (Zhang et al., 2015), the Himalaya (Gardelle et al., 2011; Nie et al., 2017) (Gardelle 89 et al., 2011; Nie et al., 2017), the HMA (Wang et al., 2020; Chen et al., 2021) (Chen et al., 90 2021; Wang et al., 2020), the Tien Shan (Wang et al., 2013) (Wang et al., 2013), the Alaska 91 92 (Rick et al., 2022), the Greenland (How et al., 2021) and the northern Pakistan (Ashraf et al., 93 2017) (Ashraf et al., 2017). However, the latest glacial lake mapping in 2020 is still absent 94 along the CPEC. Among existing studies, Landsat archival images are the most widely used 95 due to their multi-decadal record of earth surface observations, reasonably high spatial 96 resolution (30 m), and publicly available distribution (Roy et al., 2014) (Roy et al., 2014). Freely available Sentinel-2 satellite images show a better potential than Landsat in glacial 97 98 lake mapping and inventories due to their higher spatial resolution (10 m) and a global 99 coverage, but have only been available since late 2015 (Williamson et al., 2018; Paul et al., 100 2020) (Williamson et al., 2018; Paul et al., 2020). Glacial lake inventories using Sentinel images are relatively scarce at regional scales, and studies of the latest glacial lake mapping 101

102 as well as comparisons of glacial lake datasets derived from Sentinel and Landsat103 observations are still lacking.

104 Discrepancies between various glacial lake inventories (Zhang et al., 2015; Shugar et al., 2020; Wang et al., 2020; Chen et al., 2021; How et al., 2021) (Zhang et al., 2015; Shugar et-105 al., 2020; Chen et al., 2021; Wang et al., 2020) result from differences in mapping methods, 106 107 minimum mapping units, definition of glacial lakes, time periods, data sources and other 108 factors. For example, manual vectorization method was widely adopted at the earlier stage for 109 its high accuracy. However, it is time-consuming associated with high labor intensity and is only practical at regional scales (Zhang et al., 2015; Wang et al., 2020) (Zhang et al., 2015; 110 111 Wang et al., 2020). Automated and semi-automated lake mapping methods, such as band-112 ratio amulti-spectral indexices classificationnd object-oriented classification (Gardelle et al., 2011; Nie et al., 2017; Zhang et al., 2018; How et al., 2021) (Gardelle et al., 2011; Zhang et 113 al., 2018; Nie et al., 2017), have been developed to improve the efficiency of glacial lake 114 inventories using optical images, although artificial manual modification is often unavoidable 115 to assure the quality of lake data impacted by cloud cover-in optical images, mountain 116 117 shadows, seasonal snow cover and frozen lake surfaces (Sheng et al., 2016; Wang et al., 2017, 2018) (Sheng et al., 2016; Wang et al., 2017; Wang et al., 2018). SBackscatter images from 118 119 Synthetic aAperture #Radar (SAR)backscatter classification (Wangchuk and Bolch, 2020; 120 How et al., 2021) wasere used to remove the impact of cloud cover for lake mapping. Besides, 121 other approaches such as hydrological sink detection using DEM (How et al., 2021) orand land surface temperature-based detection method (Zhao et al., 2020) were also used for lake 122 123 inventories. Different classification methods impact the results of lake mapping and 124 monitoring. data, such as sentinel-1, are also increasingly used in ice lake extraction due totheir undisturbed nature by clouds (Zhang et al., 2020; Wangchuk and Bolch, 2020). In-125 126 addition, some scholars also monitor lakes based on their temperature and other-127 characteristics (Zhao et al., 2020). 128 Type Dam type –classification of glacial lakes provides a crucial attribute for glacier-lake 129 interactions and risk assessment (Emmer and Cuřín, 2021) (Emmer and Cuřín, 2021). So far, we are lacking there has not been a unified standard for the classification system of glacial 130 lakes (Yao et al., 2018) (Yao et al., 2018). Existing classification systems are mainly for their 131 132 respective research purposes, mainly based on the relative positions of glacial lakes and 133 glaciers, the supply conditions of glaciers, and the attributes of dams. In addition to different 134 classification standards, the same typespecies of glacial lakes may also have different names given by different scholars. For example, ice-marginal (Carrivick and Quincey, 2014; 135 136 Carrivick et al., 2020), ice-contact (Carrivick and Tweed, 2013) and proglacial (Nie et al., 137 2017) lakes all represent glacial lakes sharing the boundary with glaciers. Glacier lakes in 138 currently available datasets have been traditionally categorized by their spatial relationship 139 with upstream glaciers (Gardelle et al., 2011; Wang et al., 2020; Chen et al., 2021) (Gardelleet al., 2011; Chen et al., 2021; Wang et al., 2020), and classification attributes considering the 140 141 formation mechanism and the properties of dams are rare or incomplete in the CPEC (Yao et al., 2018; Li et al., 2020) (Li et al., 2021; Yao et al., 2018). Therefore, an up-to-date glacial 142 lake dataset with critical, quality-assured parameters (e.g. lake types) is necessary. 143 144 This study aims to (1) employ both Landsat 8 and Sentinel-2 images to create an up-to-date glacial lake dataset in the CPEC to accurately document its detailed lake distribution in 2020; 145

- 146 (2) reveal glacial lake changes and the spatial heterogeneity across mountains and basins in
- 147 the CPEC using consistent 30-m Landsat images at three time periods (1990, 2000 and 2020);
- 148 and (3) share the glacial lake inventories with a range of critical attributes to benefit
- 149 hazardous risk assessment of GLOFs and glacio-hydrological modeling in the HMA.

150 **2 Study area**

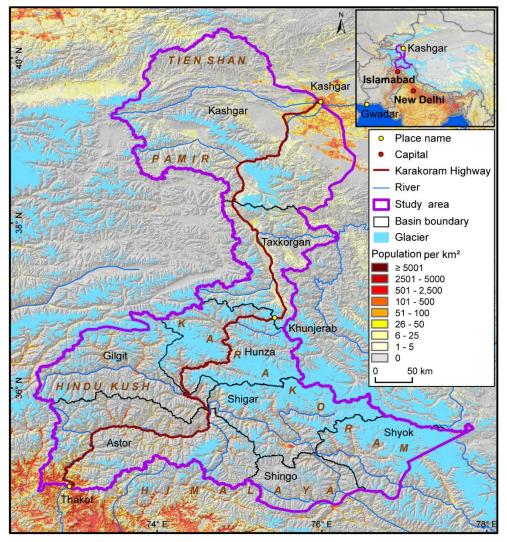


Figure 14. Location of the study area and distribution of glaciers, mountains, basins and population.

The study area (Figure 1Figure 1) covers all the drainage basins along Karakoram Highway starting from Kashgar and ending at Thakot, with a total area of $\sim 125,000 \text{ km}^2$. The upper Indus basins beyond the Pakistani-administrated border are excluded in this study due to little impact of GLOFs there on CPEC infrastructures. The entire study area is divided into eight 157 sub-basins, covering most of the Karakoram with the highest altitude up to 8611 m, western 158 159 Himalaya and Tien Shan, eastern Hindu Kush and Pamir mountains. The 9710 glaciers in the study area cover a total area of 17,447 km^2 and nearly 60% of glaciers are distributed in the 160 Karakoram (5818 glaciers with a total area of 14,067.52 km²) (RGI Consortium, 2017) (RGI 161 Consortium, 2017). Most glaciers in the western Himalaya and eastern Hindu Kush are losing 162 mass in the context of climate change (K ääb et al., 2012; Yao et al., 2012; Brun et al., 2017; 163

164 Shean et al., 2020; Hugonnet et al., 2021) (K äb et al., 2012; Yao et al., 2012; Shean et al., 2020; Brun et al., 2017; Hugonnet et al., 2021), whereas the glaciers in the eastern 165 Karakoram and Pamir have shown unusually little changes, including unchanged, retreated, 166 advanced and surged glaciers (Hewitt, 2005; K ääb et al., 2012; Bolch et al., 2017; Brun et al., 167 2017; Shean et al., 2020; Nie et al., 2021) (Nie et al., 2021; Brun et al., 2017; Shean et al., 168 169 2020; K ääb et al., 2012; Hewitt, 2005; Bolch et al., 2017). The spatially heterogeneous 170 distribution and changes of glaciers are primarily explained as a result of differences in the dominant precipitation-bearing atmospheric circulation patterns that include the winter 171 172 westerlies the Indian summer monsoon, their changing trends and their interactions with local extreme topography (Yao et al., 2012; Azam et al., 2021; Nie et al., 2021) (Azam et al., 2021; 173 174 Nie et al., 2021; Yao et al., 2012).

175 **3 Data sources**

Both Landsat and Sentinel images have been employed to map glacial lakes between 1990 176 and 2020 in the CPEC (Figure 2Figure 2). A total number of 98-71 Landsat Thematic Mapper 177 (TM), Thematic Mapper Plus (ETM+) and Landsat 8 Operational Land Imager (OLI) images 178 179 with a consistent spatial resolution of 30 m were downloaded from the United States 180 Geological Survey Global Visualization Viewer (GloVis, https://glovis.usgs.gov/app/) to be 181 used to create glacial lake inventories in 1990, 2000 and 2020. High-quality Landsat images around 2010 are insufficient to cover the entire study area, so we had to give up glacial lake 182 183 mapping in 2010 as a result of Landsat 7's scan-line corrector errors and significant cloud 184 covers. In addition, 40-39 Sentinel-2 images were downloaded from Copernicus Open Access 185 Hub (https://scihub.copernicus.eu/) to produce the 10-m resolution glacial lake inventory in 2020. 186

187 Cloud and snow covers heavily affect the usability of optical satellite images (Wulder et al., 2019) (Wulder et al., 2019) and their availability in the entire study area, so we took 188 189 advantage of the images acquired before and after each of the baseline years 1990, 2000 and 190 2020 to construct the glacial lake inventories. To minimize the impact of intra-annual changes of glacial lakes, most of used images (8582% for Sentinel and 8275% for Landsat) were 191 192 acquired from August to October in the given baseline year with cloud coverage of <20% for 193 each image. For some specific scenes where cloud cover exceeded the threshold of 20%, we 194 selected more than one image to remedy the effect of cloud contamination (Nie et al., 2010, 195 2017; Jiang et al., 2018) (Nie et al., 2010; Nie et al., 2017; Jiang et al., 2018).

196 Other datasets used include the Randolph Glacier Inventory version 6.0 (Pfeffer et al., 2014; RGI Consortium, 2017) (Pfeffer et al., 2014; RGI Consortium, 2017) and the Glacier 197 198 Area Mapping for Discharge from the Asian Mountains (GAMDAM) glacier inventory 199 (Sakai, 2019) (Sakai, 2019). These two glacier datasets were used to determine glacial lake 200 attributestypes, such as ice-contact, ice-dammed and unconnected-glacier-fed lakes. The 201 Shuttle Radar Topography Mission Digital Elevation Model (SRTM DEM) at a 1-arc second 202 (30 m) resolution (Jarvis et al., 2008) (Jarvis et al., 2008) was employed to extract the 203 altitudinal characteristics of the glacial lakes. The absolute vertical accuracy of the SRTM 204 DEM is 16 m (90%) (Rabus et al., 2003; Farr et al., 2007) (Farr et al., 2007; Rabus et al., 205 $\frac{2003}{1000}$. We also applied other published glacial lake datasets for comparative analysis. They

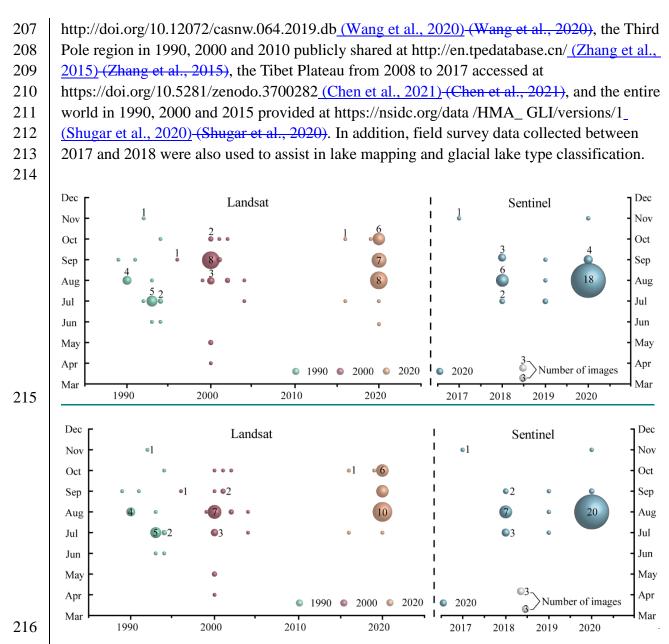


Figure 22. Acquisition years and months of Landsat and Sentinel images selected for glacial lake inventories. The bubble size indicates the available image number.

219 **4 Glacial lake inventory methods**

220 **4.1 Definition of glacial lakes**

217

218

221 We consider a glacial lake as one that formed as a result of modern or ancient glaciation. 222 Contemporary glacial lakes are easily recognized using a combination of glacier inventories 223 and remote sensing images. Ancient glacial lakes can be identified from periglacial 224 geomorphological characteristics, including moraine remnants and U-shaped valleys that are 225 discernible from satellite observations (Post and Mayo, 1971; Westoby et al., 2014; Nie et al., 2018; Mart n et al., 2021) (Post and Mayo, 1971; Nie et al., 2018; Mart n et al., 2021; 226 Westoby et al., 2014). Landslide-dammed lakes (Chen et al., 2017) (Chen et al., 2017) in the 227 228 periglacial environment were excluded in our inventories because of their irrelevance to

229 glaciation. We abandoned the definition that considers all lakes surrounding a specific 230 buffering distance of other glaciers also as glacier lakes, although this definition has been 231 widely used in previous studies assuming glacial meltwater as the main water supply (Zhang 232 et al., 2015; Wang et al., 2020) (Zhang et al., 2015; Wang et al., 2020). This is because the 233 contribution of glacial meltwater to the lake supply is arduous to be quantified without an 234 accurate modeling of the cryosphere-hydrological processes (Lutz et al., 2014) (Lutz et al., 235 2014). All glacial lakes in the study area were mapped according to our definition without 236 any regard to buffering distance limit between lakes and of glaciers. We were able to 237 implement this definition by carefully leveraging the spectral properties of glacial lakes and 238 the periglacial geomorphological features that are often evident in remote sensing images (see 239 more in sections 4.3 and 4.4).

240 4.2 Interactive lake mapping

246

247

248

249

250

251

252 253

254

255

256

257

258

A human-interactive and semi-automated lake mapping method (Wang et al., 2014; Nie et al., 2014; Nie et al., 2017, 2020) (Wang et al., 2014; Nie et al., 2017; Nie et al., 2020) was adopted to accurately
extract glacial lake extents using Landsat and Sentinel-2 images, based on the Normalized
Difference Water Index (NDWI) (Mcfeeters, 1996) (Mcfeeters, 1996). The NDWI uses the
green and near infrared bands and is calculated by the following equation:

$$NDWI = \frac{Band_{Green} - Band_{NIR}}{Band_{Green} + Band_{NIR}}$$
(1)

where the green band and near infrared band were provided by both Landsat and Sentinel multispectral images.

Specifically, <u>the method calculated thethe NDWI histogram</u>the method automaticallygenerated the histogram based on the pixels withof NDWI in each user-defined <u>and</u> <u>manually-drawnew</u> region of interest. The NDWI threshold that separates lake surface from land was interactively determined by screening the NDWI histogram against the lake region in the imagery (Wang et al., 2014; Nie et al., 2020) (Nie et al., 2020; Wang et al., 2012). This way, the determined NDWI threshold can be well-tuned to adapt various spectral conditions of the studied glacier lakes. The raster lake extents segmented by the thresholds were then <u>automatically</u> converted to vector polygons. We first completed the glacial lake inventory in 2020 using this interactive mapping method, and the 2020 inventory was then used as a reference to facilitate the lake mapping for other periods.

259 The minimum mapping unit (MMU) was set to 5 pixels for both Landsat (0.0045 km^2) and Sentinel-2 images (0.0005 km^2) in this study. MMU determines the total number and area of 260 261 glacial lakes in the dataset, and varies in the previous studies, such as 3 pixels (Zhang et al., 262 2015) (Zhang et al., 2015), 9 pixels (Chen et al., 2021) (Chen et al., 2021), or 55 pixels (Shugar et al., 2020) (Shugar et al., 2020) for Landsat images for various objectives and 263 264 spatial scales. While a smaller threshold leads to a large quantity of lakes mapped, it also 265 generates larger mapping noises or uncertainties. Considering this signal-noise balance and our focus on identifying prominent glacier lake dynamics in the study area, we opted to use 5 266 267 pixels as the <u>minimum mapping unitMMU</u> for both Landsat and Sentinel-2 images. Several procedures were taken to assure the quality assurance and quality control for lake

Several procedures were taken to assure the quality assurance and quality control for lake
mapping, including 1) visual inspection and modification for each lake based on Landsat,
Sentinel-2 and Google Earth high-resolution images overlaying preliminarily lake boundary

271 extraction at the given time period; 2) time series check for Landsat-derived glacial lake 272 datasets from 1990 and 2020, and cross-check between Landsat and Sentinel-2-derived lake 273 dataset in 2020 to reduce errors of omission and commission; 3) topological validation of 274 glacial lake mapping, such as repeated removal, elimination of small sliver polygons; and 4) 275 logical check for lake types between two classification systems of glacial lakes. False lake 276 extents resulting from cloud or snow cover, lake ice, and topographic shadows (Nie et al., 277 2017, 2020) (Nie et al., 2020; Nie et al., 2017) and were modified using alternative images 278 acquired in adjacent years. Those procedures were time-consuming, but helped to minimize 279 the effect of cloud and snow covers, lake mapping errors, and to maximize the quality of the 280 produced lake product and the derived glacial lake changes.

281 4.3 Classification of glacial lakes

300

301

282 Two glacial lake classification systems (GLCS) have been established based on relationship 283 of interaction between glacial lakes and glaciers as well as lake formation mechanism and 284 dam material properties. In the first GLCS (GLCS1), glacial lakes were classified into four 285 types based on their spatial relationship to upstream glaciers: supraglacial, proglacialice-contact, unconnected-glacier-fed lakes, and non-glacier-fed lakes according to 286 287 Gardelle et al. (2011) (2011) and Carrivick et al. (2013). Alternatively, combining the formation mechanism of glacial lakes and the properties of natural dam features, glacial lakes 288 289 were classified into five categories (herein named GLCS2) modified from Yao's classification 290 system (2018)-(2018): supraglacial, end-moraine-dammed, lateral-moraine-dammed, 291 glacial-erosion lakes and ice-blocked-dammed lakes. Subglacial lakes were excluded due to 292 the mapping challengelimitation in spectral properties of optical from spectral satellite 293 images alone. Characterization and examples for each type are provided in Table 1 Table 1 and 294 Table 2Table 2. Individual glacial lakes were categorized to the specific types for each GLCS 295 according to available glacier inventory data, geomorphological and spectral characteristics 296 interpreted from Landsat, Sentinel and Google Earth images. The synergy of these two GLCSs is beneficial to predicting glacier-lake evolutions and providing fundamental data for 297 298 glacial lake disaster risk assessment. 299

Table 11. Classification system of glacial lake types according to the relationship between glacial lakes and glaciers (© Google Earth 2019).

Lake types	Characteristics	Landsat	Sentinel	Google earth
1 0	Lakes formed on the surface of glaciers, generally dammed by ice and thin debris. Case location: 35 43'49.74" N 76 93'53.88" E			

Proglacial <u>Ice</u> - contact	Lakes dammed by moraine, ice or bedrock, supplied by glacial meltwater and <u>shared</u> <u>boundaryconnected</u> with glaciers. Case location: 39 '09'32.40" N 73 °43'12.00" E		
Unconnected- glacier-fed	Lakes currently supplied by upstream glacial meltwater but disconnected with glaciers. Case location: 35 '47'60.00" N 72 '55'15.60" E		
Non-glacier-f ed	Lakes formed by glaciology, dammed by moraine or bed rock, and currently not supplied by glacial meltwater. Case location: 34 '50'39.99" N 74 '48'29.31" E		

Table 22. Classification system of glacial lake types according to the formation mechanism of glacial lakes and dam material properties (© Google Earth 2019).

Lake types	Characteristics	Landsat	Sentinel	Google earth
Supraglacial	Lakes formed on the surface of glaciers, generally dammed by ice and thin debris. Case location: 36 467.39" N 74 207.59" E			
End-moraine <u></u> damme	dLakes formed behind moraines as a result of glacier retreat and downwasting. Case location: 35 °42'50.40" N 73 °09'57.60" E			
Lateral- _morainedammed	Lakes formed behind lateral glacial moraine ridges and dammed by debris, different from ice-blockedice-dammed glacial lake. Case location: 38 '28'45.62" N 75 '20'52.30" E			
Glacialerosion	Lakes formed in depressions created by glacial over-deepening. Bedrock dam dominates, partially superimposed by top moraine <u>`</u> in rugged terrain. Dams are unclear in the satellite images. Case location: 35 '55'55.56'' N 73 '38'20.13'' E			
Ice-blocked dammed	Lakes formed behind glaciers, dammed by glacier ices (partially covered by debris on the top). Case location: 35 '28'31.32'' N 77 '30'46.81'' E			

305

306 4.4 Attributes of glacial lake data

A total of 17 attribute fields were input into our glacial lake datasets (-<u>Table 3</u><u>Table 3</u>). They
include lake location (longitude and latitude), lake elevation (centroid elevation), orbital
number of the image source, image acquisition date, lake area, lake perimeter, lake types of
the two GLCSs, mapping uncertainty, and the country, sub-basin, and mountain range
associated with the lake. Amongst the attributes, lake location was calculated based on the
centroid of each glacial lake polygon associated with the DEM, N represents northing and E

- 313 represents easting. Orbital number of the image source was filled with the corresponding
- satellite image, with the codes expressed as "PxxxRxxx" or "Txxxxx", where P and R 314
- 315 indicate the path and row for Landsat image and T represents the tile of Sentinel image
- associated with 5 digits code of military grid reference system. Area and perimeter were 316
- automatically calculated based on glacial lake extents. Lake types were attributed using the 317
- 318 characterization and interpretation marks described in Section 4.3. Mapping uncertainty was
- 319 estimated using our modified equation which will be introduced in section 4.5 and
- 320 supplementary appendix tutorial. Located country, sub-basin and mountain range of each
- glacial lake was identified by overlapping the geographic boundaries of countries, basins and 321 322 mountain ranges.
- 323
- Table 33. Classification system of glacial lake types according to the formation mechanism of glacial lakes
 324 and dam material properties.

Field Name	Туре	Description	Note
FID or	Object ID	Unique code of glacial lake	Number
OBJECTID			
Shape	Geometry	Feature type of glacial lake	Polygon
Latitude	String	Latitude of the centroid of glacial lake	Degree minute second
		polygon	
Longitude	String	Longitude of the centroid of glacial lake	Degree minute second
		polygon	
Elevation	Double	Altitude of the centroid of glacial lake	Unit: meter above sea level
		polygon	
IMGSOURCE	String	Path and row numbers for Landsat image	PxxxRxxx or Txxxxx
		based on World Reference System 2 or Tile	
		number for Sentinel image based on military	
		grid reference system	
ACQDATE	String	Acquisition date of source image	YYYYMMDD
GLCS1	String	The first classification system of glacial lakes	Supraglacial,
		based on relationship of interaction between	ProglacialIce-contact,
		glacial lakes and glaciers	Unconnected-glacier-fed,
			None-glacier-fed
GLCS2	String	The second classification system of glacial	Supraglacial,
		lakes based on lake formation mechanism and	End-moraine-dammed,

Field Name	Туре	Description	Note
		dam material properties	Lateral-moraine-dammed,
			Glacial-
			erosionGlacial-erosion and
			Ice-blockedIce-dammed
Basin	String	Basin name where glacial lake locates in	
Mountains	String	Mountain name where glacial lake locates in	
Country	String	Country name where glacial lake locates in	
Perimeter	Double	Perimeter of glacial lake boundary	Unit: meter
Area	Double	Area of glacial lake coverage	Unit: square meter
Uncertainty	Double	Uncertainty of glacial lake mapping estimated	Unit: square meter
		based on modified Hanshaw's equation	
		(2014).	
Operator	String	Operator of glacial lake dataset	Muchu, Lesi
Examiner	String	Examiner of glacial lake dataset	Yong, Nie

326 4.5 Improved uncertainty estimating method

327 We modified Hanshaw's (2014) (2014) equation that had been used to calculate lake-area 328 mapping uncertainty. Lake perimeter and displacement error are widely used to estimate the 329 uncertainty of glacier and lake mapping from satellite observation (Carrivick and Quincey, 330 2014; Hanshaw and Bookhagen, 2014; Wang et al., 2020) (Wang et al., 2020; Li et al., 2020; Zhang et al., 2015; Gardelle et al., 2011; Carrivick and Quincey et al., 2014). Hanshaw and 331 332 Bookhagen (2014) (2014) proposed an equation to calculate the error of area measurement by the number of edge pixels of the lake boundary multiplied by half of a single pixel area. The 333 334 number of edge pixels is simply calculated by the perimeter divided by the grid size. The equation is expressed as below: 335

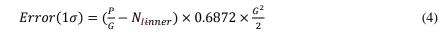
336

$$Error(1\sigma) = \frac{P}{G} \times 0.6872 \times \frac{G^2}{2}$$
(2)

337

$$D = \frac{Error(1\sigma)}{4} \times 100\% \tag{3}$$

Where *G* is the cell size of the remote sensing imagery (10 m for Sentinel-2 image and 30 m
for Landsat image). *P* is the perimeter of individual glacial lake (m), and the revised
coefficient of 0.6872 was chosen assuming that area measurement errors follow a Gaussian
distribution. Relative error (*D*) was calculated by equation 3, in which A is the area of an
individual glacial lake.



Where N_{linner} is the number of inner nodes (inflection points) of each lake. The modified equation is also suitable for lakes with islands (as illustrated in Figure 3b).

For polygons without islands (Figure \$3a), use the following equation:

$$N_{linner} = \left(\frac{N_{Total} - 4 - 1}{2}\right) \tag{5}$$

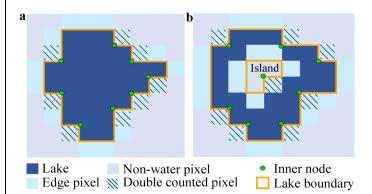
(6)

 N_{Total} is the total number of nodes, including both the outer and inner. N_{Total} were calculated by the "Field Calculator" in ArcGIS, in some cases, it is necessary to remove the redundant nodes before calculating the total number of nodes (See the Supplement for more details). An inner node is a polygon vertex where the interior angle surrounding it is greater than 180 degrees. An outer node is the opposite of the inner node, where the interior angle is less than 180 degrees. We found that the outer nodes are usually four more than the inner nodes in our glacial lake dataset. The total nodes in ArcGIS contain one overlapping node to close the polygon, meaning the endpoint is also the startpoint. This extra count was deleted in the calculation (equation 5).

 $N_{li=nner} = \left(\frac{N_{Total} - (N_{li=sland} + 1) \times 5}{2}\right)$

 $N_{lisland}$ is the number of islands within each polygon. A calculation method of $N_{lisland}$ is

For polygons with island (Figure \$3b) use the following equation:



given in the Supplement.

368 369

Figure 33. Sketch of estimating the actual edge pixels for uncertainty calculation of individual glacial lake (with and without islands).

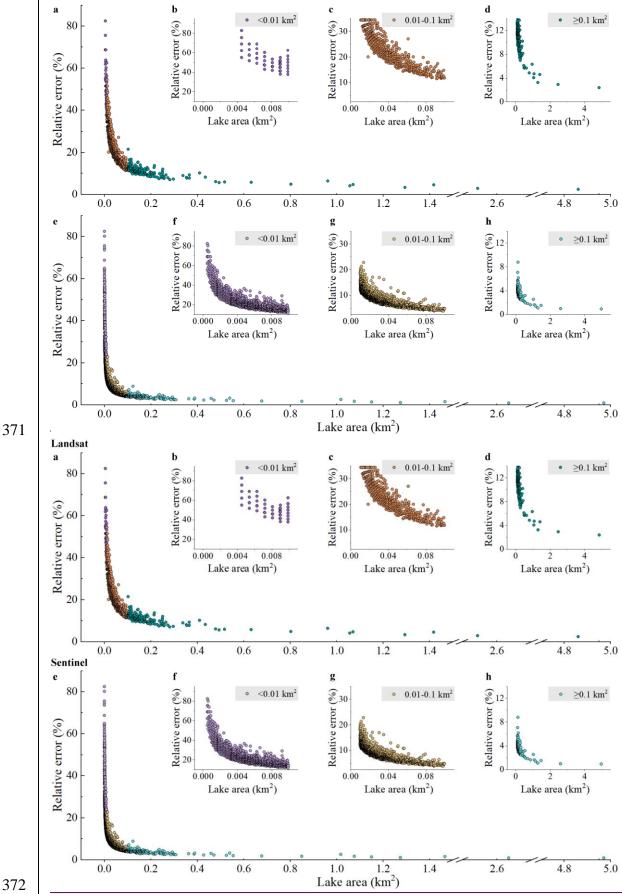




Figure 44. Relationships between individual lake size and its estimated relative error for glacial lakes of all

or specific size ranges in study area. Error estimation is based on the modified equation and lake data
extracted from Landsat (a-d) and Sentinel images (e-h).

377 The uncertainty estimated from our improved equation shows that the relative error of individual glacial lake decreases when lake size increases or cell size of remote sensing 378 379 images reduces (Lyons et al., 2013; Carrivick and Quincey, 2014)-(Lyons et al., 2013; 380 Carrivick and Quincey, 2014) (Lyons et al., 2013) (Figure 4Figure 4). Total area error of glacial lakes in study area is approximate ±14.98 km² and ±8.45 km² in 2020 for Landsat and 381 Sentinel images, respectively, and the average relative error is $\pm 17.36\%$ and $\pm 8.15\%$. 382 Generally, small lakes have greater relative errors. For example, the mean relative error is 383 35.38% for Landsat derived glacial lakes between 0.0045 and 0.1 km² and 10.63% for glacial 384 lakes greater than 0.1 km². The mean area error of Sentinel-derived glacial lakes is almost 385 one sixth of that extracted from Landsat images for glacial lakes of all or specific size group. 386

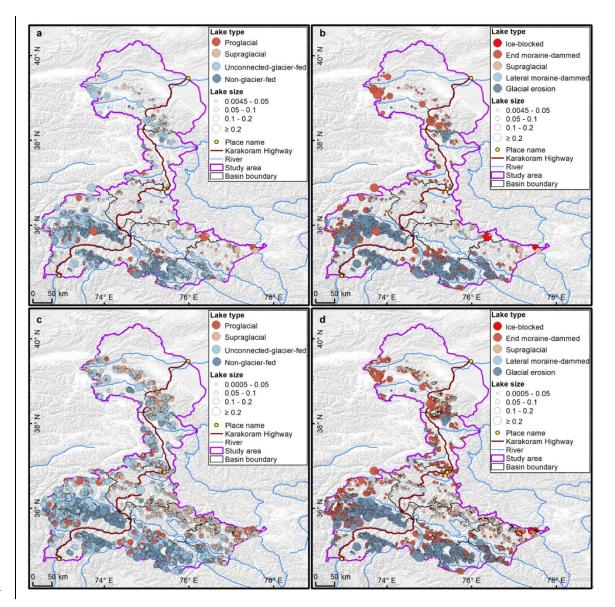
5 Results

376

388 5.1 Glacier lake distribution and changes observed from Landsat

We mapped 2,234 glacial lakes for 2020 across the studied CPEC from Landsat-8 images, 389 with a total area of $86.31 \pm 14.98 \text{ km}^2$ (Figure 5Figure 5 a and b). The majority of these glacial 390 lakes (1,870 or 83.71%) are smaller than 0.05 km² and contribute 36.5% of the total area. 45 391 (2.01%) of the lakes are larger than 0.2 km^2 and contribute 28.8% of the total area (Figure 392 6Figure 6). With the increase of lake size, the abundance (count) of glacial lakes consistently 393 394 decreases but the total lake area first reduces and then increases. Unconnected-glacier-fed 395 lakes are dominant in the first classification system, followed by non-glacier-fed lakes 396 (Figure 7Figure 7) whereas glacial erosion glacial-erosion lakes dominate at both number (1478) and area (57.02 km²) in the second classification system (Figure 8Figure 8), followed 397 398 by end moraine-dammedend-moraine-dammed lakes and supraglacial lakes. Among the 399 classified lakes, 137 are proglacialice-contact lakes and cover an area of 5.56 km², implying a 400 higher mean size of proglacier-ice-contact lakes than supraglacial lakes.

401 Glacial lakes are spatially heterogeneous among various mountain ranges and basins in the study area. Himalaya sub-region has the maximum glacier lake count and area across the 402 403 entire study area, followed by Hindu Kush. Supraglacial lakes are mainly distributed in the 404 Karakoram but they cover less area than those in the Pamir. Tien Shan has fewer glacial lakes. 405 Astor, Gilgit and Shingo basins have the largest percentages of glacier lakes in both number and area (>17%) (Figure 9Figure 9a), and each of the other basins contributes less than 10% 406 except Kashgar basin in area due to several large ancient glacial lakes. Glacial lakes of less 407 than 0.05 km² dominate in number within each basin and the total number decreases as lake 408 size increases. Small lakes consistently account for the maximum percentage in area except 409 410 Kashgar basin as a result of the disproportionally large lakes.



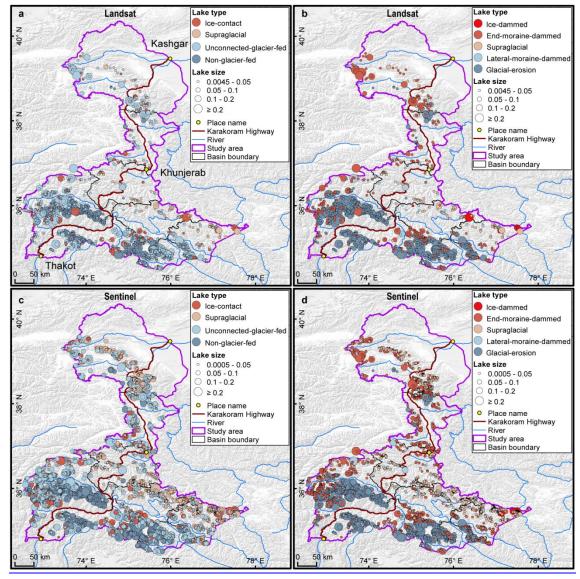




Figure 55. Distribution of glacial lakes in 2020 extracted from Landsat (a, b) and Sentinel (c, d) images.Panels a and b-c are classified by GLCS1, and GLCS2 for sub-graph e-b and d.

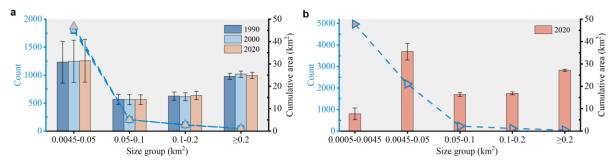


Figure 66. Statistics of different sizes of glacial lakes in the study area from 1990 to 2020. Panels a and b were derived from Landsat and Sentinel images, respectively.

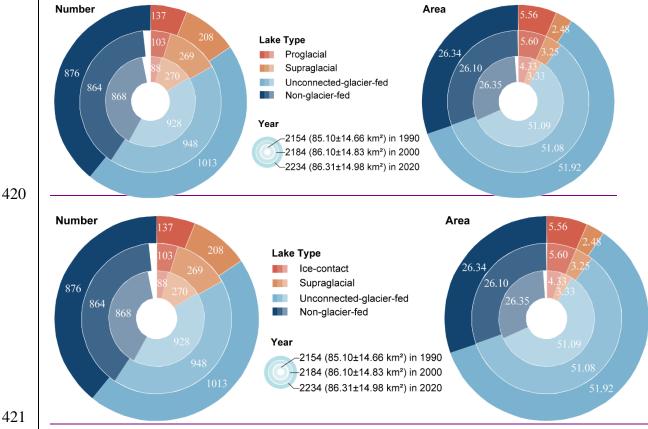
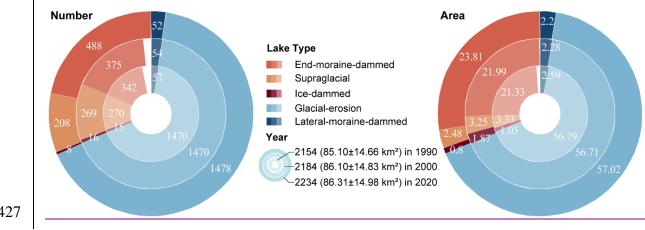


Figure 77. Number and area of different types of glacial lakes classified based on the condition of glacier supply in the study area. The outermost ring represents glacial lake data in 2020, middle ring for 2000 and innermost ring for 1990. Lake number and area in 2020 were selected as reference, meaning a concept of "100 %" for a complete ring. Labeled values are scaled in degrees rather the radius of rings.



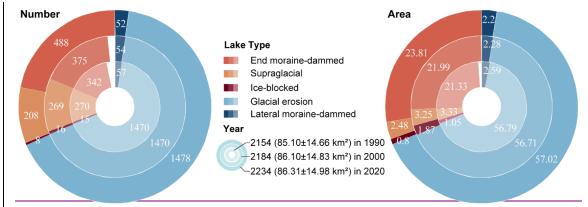


Figure §8. Number and area of different types of glacial lakes classified based on glaciation and nature of dam in the study area. The outermost ring represents glacial lake data in 2020, middle ring for 2000 and innermost ring for 1990. Lake number and area in 2020 were selected as reference, meaning a concept of "100 %" for a complete ring. Labeled values are scaled in degrees rather the radius of rings.

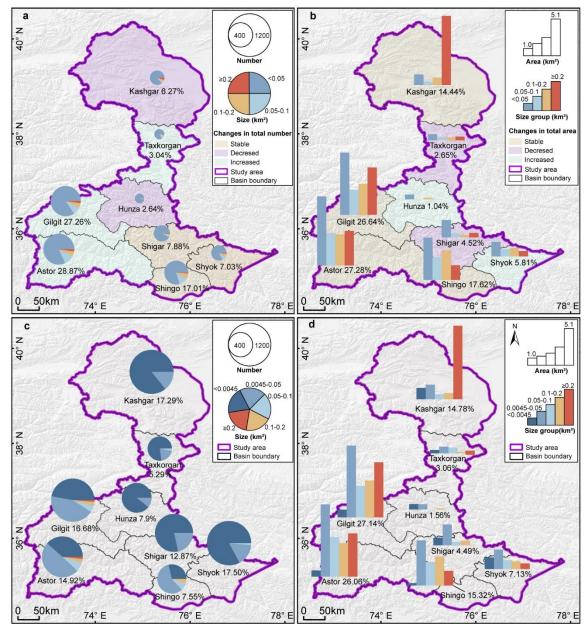


Figure 99. Distributions and changes in count and area of glacial lakes. Percent of glacial lakes in number or area is labeled in each basin. Pie charts present the number of glacial lakes at various size groups between basins (a and c) and bar charts represent total area of glacial lakes at different size groups in each basin (b and d). The background colors represent changes in total number and area between 1990 and 2020 based on Landsat derived dataset (a and b) and distribution of Sentinel derived glacial lakes in 2020 among basins are shown in sub-graphs c and d.

The total number and area of glacial lakes in the study remain relatively stable with a slight increase between 1990 and 2020, and the changes in count and area among various types of glacial lakes vary substantially (Figure 7Figure 7 and Figure 8Figure 8). From 1990 to 2020, the total number of glacial lakes increased by 80 or 3.70%, while the area grew by 1.21 km^2 (or 1.42%) while the area grew by a less extent (1.21 km^2 or 1.42%). Small lakes (<0.05 km²) continuously increased in number and area, and contributed most in the total lake expansion (Figure 6Figure 6). Lakes in the size group of 0.05-0.1 km² remained stable. The total area of

450 lakes greater than 0.1 km^2 consistently increased.

In the GLCS1, unconnected-glacier-fed lakes have the largest increase in number, followed by proglacialice-contact and non-glacier-fed lakes, whereas supraglacial lakes decreased by 62 in count. ProglacialIce-contact lakes expanded by 1.24 km² (equaling an increase of 26% in proglacialice-contact lakes), contributed one third of the total area increase. Supraglacial lakes decreased by 0.85 km² in area whereas the areas of unconnected-glacier-fed and non-glacier-fed lakes remained stable as a result of disconnections from glaciers (Figure 457 <u>7Figure 7</u>).

In the GLCS2, end moraine dammedend-moraine-dammed lakes increased by 2.48 km²
and contributed most of the glacier lake area expansion, whereas supraglacial,
ice blockedice-dammed and lateral-moraine-dammed lakes decreased slightly in both
number and area. Glacial erosionGlacial-erosion lakes accounted for the maximum
percentage (about 66% for both count and area) in each time period and remained stable
(Figure 8Figure 8).

Spatially, glacial lake changes in number and area vary among different mountain ranges 464 and basins between 1990 and 2020 in the study area. Glacial lakes across the west Himalaya 465 466 and Hindu Kush increased both in number and area between 1990 and 2020 whereas the total 467 number of glacial lakes decreased in the Karakoram, Pamir and Tien Shan of study area (Table 4Table 4). The total area of glacial lakes continued to increase in the Hindu Kush, but 468 469 decreased between 1990 and 2000 and increased between 2000 and 2020 in the Himalaya. 470 The total number of glacial lakes continuously decreased in the Pamir and Tien Shan in the 471 past three decades but increased at the first stage and decreased after in the Karakoram. The 472 total area of glacial lakes persistently grew in the Pamir whereas fluctuated in the Tien Shan 473 and Karakoram.

The total numbers of glacial lakes in Shingo, Shigar and Shyok basins were stable (Figure 9Figure 9a and b); however, the areal changes were less so, including being a stable trend for Shingo, decreasing for Shigar, and increasing for Shyok. The total number of glacial lakes increased in the basins of Astor, Gilgit and Taxkorgan, whereas the total area of glacial lakes remained stable in Astor and Gilgit basins and decreased in Taxkorgan basin. The total numbers of lakes in Kashgar and Hunza basins decreased, whereas the total area of glacial lakes remained stable in Kashgar and increased in the Hunza basin.

481 482

Table <u>44</u>. Distributions in count and area (km²) of glacial lakes among mountain ranges within the study area.

		~ /	0	U	U	5
Source and year	Tien Shan	Karakoram	Pamir	Hindu Kush	Himalaya	Total
Landsat in 1990	10 (0.12)	370 (11.11)	178 (13.73)	780 (28.33)	816 (31.81)	2154 (85.10)
Landsat in 2000	7 (0.11)	393 (11.76)	163 (13.96)	792 (28.50)	829 (31.77)	2184 (86.10)
Landsat in 2020	5 (0.17)	334 (10.10)	182 (14.14)	835 (29.25)	878 (32.65)	2234 (86.31)
Sentinel in 2020*	11 (0.21)	479 (11.69)	262 (15.71)	880 (34.96)	959 (33.39)	2591 (95.96)

483 *Note: Glacial lake greater than 4500 m² are calculated for Sentinel-2 derived dataset in order to be in line with Landsat
484 derived dataset.

485 5.2 Glacier lake distribution observed from Sentinel-2

486 Sentinel-derived results shows that there are 7,560 glacial lakes $(103.70 \pm 8.45 \text{ km}^2)$ in 2020 487 across the entire CPEC (<u>Table 5</u>) under a minimum mapping unit of 5 pixels (500 m²). 488 Similar to the pattern from Landsat mapping, the lake abundance extracted from Sentinel images is inversely related to lake size (following a typical Pareto distribution). The smallest 489 size class (0.0005-0.0045 km²) contains the maximum lake count (4,969) but the least lake 490 area $(7.73 \pm 2.62 \text{ km}^2)$ (Table 5 Table 5), which is not available in the Landsat-derived lake data 491 due to a coarser spatial resolution. In each size class, there are also a higher number of larger 492 493 glacial lakes from Sentinel than that from Landsat images. The discrepancy is mainly 494 attributed to the inconsistency of spatial resolutions and image acquisition dates and spatial 495 resolutions.

496 497

498

Table 55. Count and area of glacial lakes mapped from Sentinel and Landsat images in 2020 between
various size classes

Lake size	Glacial lakes from Sentinel	Glacial lakes from Landsat	Overlap
km ²	count (km ²)	count (km ²)	% (%)
0.0005-0.0045	4969 (7.73±2.62)		
0.0045-0.05	2182 (35.52±3.72)	1870 (31.47±9.57)	85.70 (88.60)
0.05-0.1	237 (16.37±0.89)	204 (14.07±2.18)	86.08 (85.95)
0.1-0.2	122 (16.88±0.68)	115 (15.91±1.83)	94.26 (94.25)
≥0.2	50 (27.20±0.54)	45 (24.86±1.40)	90.00 (91.40)
Total	7560 (103.70±8.45)	2234 (86.31±14.98)	

499

500 Compared with our Landsat-based product, glacial lakes from Sentinel-2 have similar 501 distribution characteristics (Figure 9Figure 9c and d) among mountain ranges, basins, types 502 and altitudinal locations (Figure 10Figure 10); meanwhile, a larger quantity of glacier lakes, with 503 more accurate boundaries and a greater total lake area, were generated from Sentinel-2 images. Taking 504 altitudinal distribution for example, the number and size of glacial lakes in the study area appear follow a 505 normal distribution against elevation for both Sentinel-2 and Landsat derived products (Figure 10Figure-506 10). The elevation of all glacial lakes mapped in 2020 based on Sentinel-2 images ranged from 2500 m to 507 5750 m (a.s.l.), with 89.58% between 3600 m and 5100 m and a mean altitude of 4421 m. The peak 508 number appears between 4500 m and 4550 m whereas the maximum area emerges between 4250 m and 509 4300 m. The anomalously large area between 3600 and 3650 m shows up in Fig. 10b because of several 510 disproportionally large lakes. Although Landsat derived lakes show a similar distribution pattern to 511 Sentinel derived lakes, the lake count and area in each altitudinal band are greater in the Sentinel product 512 due to the improved spatial resolution and image quality. 513

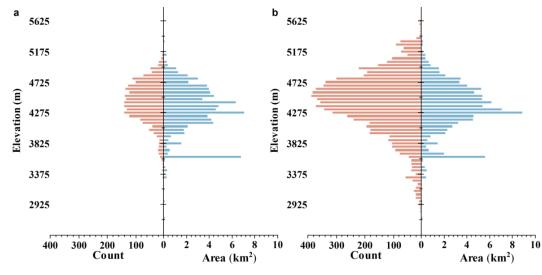
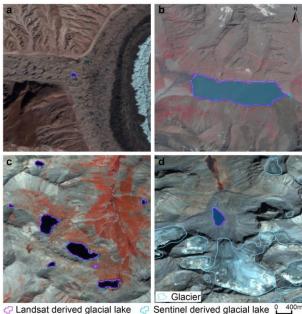


Figure 1010. Altitudinal distribution of glacial lakes in 2020 derived from Landsat (a) and Sentinel images (b)

6 Discussions

6.1 Comparison of Sentinel-2 and Landsat derived products

Glacial lakes from Landsat and Sentinel images have a high consistency in number and area with overlap rates from 85.7% to 94.26% for all lakes greater than 0.0045 km² approximately (Table 5 Table 5), implying a good potential for coordinated utility with Landsat archived observation (Figure 11Figure 11). Lake extents extracted from Landsat and Sentinel images match well for various types and sizes (Table 4Table 4). The best consistency rate reaches 94% for the glacial lakes between 0.1 km^2 and 0.2 km^2 . The difference in area of glacial lakes extracted from Landsat and Sentinel images generally lies within the uncertainty ranges. The high consistency of Sentinel-2 and Landsat derived glacial lake products in 2020 assuresimproves the value of our lake dataset. Taking the usage in assessing the assessment of GLOFs as an example, we set 0.05 km^2 as the area threshold to select object lakes, including ice-contact lakes and ice-dammed lakes that are the most active lakes and source lakes of GLOFs in the CPEC (Nie et al., 2021). A total of 24 and 29 ice-contact lakes were selected from Landsat and Sentinel-derived productsdataset, respectively. Among them, there were 4 ice-dammed lakes from the Landsat-derived productdataset, and 5 from the Sentinel-derived productdataset. All tTheese selected lakes can be used for GLOFs hazard evaluation. Because of the high consistency between our Landsat and Sentinel-based mappings, Uusers may have the flexibility are able to customizeset their own the lake size criteria to facilitate their specific purposesto select object lakes and benefit a lot from our dataset.



C Landsat derived glacial lake

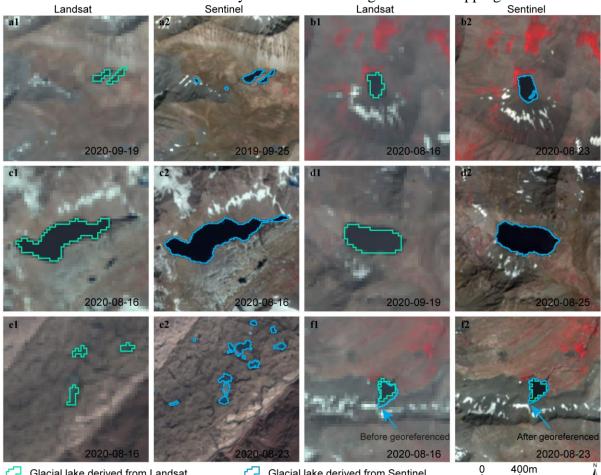
Figure 1111. High consistency of lake extents extracted from Landsat and Sentinel images. Lake types shown include supraglacial (a), glacier-fed moraine-dammed (b), unconnected glacial erosionglacial-erosion lake without glacier melt supply (c) and glacier-fed moraine-dammed (d).

Spatial resolution of satellite images plays a primary role in the discrepancies in count and area of glacial lakes extracted from Landsat (30 m) and Sentinel (10 m) observations. Due to a finer spatial resolution, Sentinel images can extract more glacial lakes and more accurate extents than those from Landsat images. We set the same 5 pixels as the minimum mapping unit for both Landsat and Sentinel images, which corresponds to a minimum area of 0.0045 km^2 and 0.0005 km^2 , respectively. The minimum mapping area results in generating nearly 5000 more lakes from Sentinel images than from Landsat images, causing the greatest discrepancy in number of the two glacial lake products (Table 5Table 5), such as Figure 12Figure 12a. Small lakes, such asfor example supraglacial lakes play an important role in understanding glaciermeltwater runoff-surface melting, and supraglacial drainage systemsglacier and lake evolutions, and cryospheric hydrology (Liu and Mayer, 2015; Miles et al., 2018). Our dataset can be used not only for GLOFs evaluation, but also for glacial lake evolution simulation and glacio-hydrological prediction. Meanwhile, Sentinel images are able to depict boundaries of glacial lake with a lower uncertainty (Figure 12Figure 12b-d). For example, some small islands and narrow channels (Figure 12Figure 12b and c) were mapped from Sentinel imagery that are-were unable to be detected in Landsat imagery.

Different acquisition dates between Sentinel and Landsat images also contribute to the discrepancy of those two glacial lake datasets. Acquiring same-day images from the two sensors were not always possible due to the impacts of cloud contaminations, topographic shadows, snow cover and revisit periods (Williamson et al., 2018; Paul et al., 2020)-(Williamson et al., 2018; Paul et al., 2020). Glacial lakes are changing temporally in the context of climate and glacier changes, taking supraglacial lakes for example that evolve dramatically in a short period (Figure 12Figure 12e). Despite our efforts of leveraging all available high-quality images, the overlap of acquisition dates between Landsat and Sentinel 568

569 images for the same location is relatively low (only 7 scenes of Sentinel images or 112 glacial 570 lakes in 2020) in this study area, and the consequential temporal gaps led to a difference in 571 the number and area of the derived glacial lakes.

Displacement between images also resulted in a certain degree of discrepancy between Landsat and Sentinel derived glacial lakes. All images used in this study have been orthorectified before download, but we still find that a fewone Sentinel images were-was not well matched with Landsat images, leading to the discrepancy between the two glacial lake datasets (Figure 12Figure 12f). We manually georeferenced the shifted images to minimize the difference between Sentinel and Landsat derived glacial lakes (Figure 12Figure 12f). Original geo-referencing accuracy is approximately approximate half of one pixel for Landsat and Sentinel image, and this displacement likely contributes a minor error to glacial lake changes at various time periods. Although we could not eliminate this intrinsic error, the error has been considered in the uncertainty assessment of our glacial lake mapping.



Glacial lake derived from Landsat Figure 1212. Discrepancy of lake extents extracted from Landsat and Sentinel images.

572 573

574

575

576

577 578

579

580

581

582 583

584

Glacial lake derived from Sentinel

585 6.2 Comparison with other datasets

586 Glacial lake datasets play a fundamental role in GLOF risk evaluation, glacier change 587 prediction, and water resource availability. An increasing number of glacier lake datasets 588 have been released over the past years, and most of them were produced from long-term Landsat archives. Glacial-Regional glacial lake datasets using Sentinel images are so-scarce. 589

590 Lack of Sentinel-derived glacial lake dataset in the study area makes it impossible to compare.
 591 that we are unable to compare our product with other existing ones in the study area. Here we
 592 selected four available glacial lake datasets to compare with our Landsat-derived dataset.

Our study provides the latest glacial lake dataset (in 2020) and the most long-term Landsat 593 observation (1990 to 2020) for this study, with a range of critical attributes including two 594 595 types of classification systems. Within the same study area, our 2020 glacial lakes appear to 596 be closest to the 2018 dataset produced by Wang et al (2020) (2020), with the highest overlap 597 of greater than 74% in both number and area (Table 6 Table 6). In Wang et al. (2020) (2020), the minimum mapping unit is 6 pixels so their dataset has a smaller lake quantity. However, 598 their dataset contains all lakes within 10 km of glacier boundaries, including many large 599 600 landslide-dammed lakes that are excluded in our glacial lake mapping. As a result, their total 601 glacier lake area is greater than ours. The overlapping rates between Wang's glacial lakes (2020) (2020) in 1990 and ours are more than 69% in both number and area. However, their 602 results show a distinct increase of glacial lakes in number and area between 1990 and 2018 603 604 (Wang et al., 2020) (Wang et al., 2020) whereas our data show a more stable change between 605 1990 and 2020. One possible reason is that manually delineating glacial lakes twice by different operators during Wang's lake mapping (2020) (2020) exacerbates the errors of 606 mapping. Another reason is that their data contains landslide-dammed lakes that fluctuate 607 608 greatly with time and expanded recently. One example is the Attabad Lake (Located at 609 36 °18'22.33"N, 74 °49'34.36"E).

610 611

Table 66. Comparison of different glacial lake datasets sourced from Landsat images in the study area.

Acquisition	Method	MMU	Count	Overlap	Reference
year (period)		m ² (pixels)	(km ²)	% (%)	
1990 (1988-1993)	Manual	5400 (6)	1720 (89.68±13.69)	69.17 (76.33)	Wang et al., 2020
1990 (1990-1999)	Automated	50000 (55)	145 (20.28)	6.27 (21.66)	Shugar et al., 2020
1990 (1989-1992)	Manual	2700 (3)	622 (51.93±10.15)	27.72 (39.94)	Zhang et al., 2015
1990 (1989-1994)	Automated & Manual	4500 (5)	2154 (85.10±14.66)	—	This study
2000 (1999-2001)	Manual	2700 (3)	724 (61.41±11.91)	31.91 (46.97)	Zhang et al., 2015
2000 (2000-2004)	Automated	50000 (55)	155 (22.35)	6.78 (23.72)	Shugar et al., 2020
2008	Automated & Manual	8100 (9)	1067 (65.45)	44.14 (53.58)	Chen et al., 2021
2000 (1996-2004)	Automated & Manual	4500 (5)	2184 (86.10±14.83)	—	This study
2018 (2017-2018)	Manual	5400 (6)	1956 (102.46±15.48)	74.57 (85.63)	Wang et al., 2020
2015 (2015-2018)	Automated	50000 (55)	148 (21.45)	6.27 (22.97)	Shugar et al., 2020
2017	Automated & Manual	8100 (9)	1063 (63.23)	45.21 (57.78)	Chen et al., 2021
2020 (2016-2020)	Automated & Manual	4500 (5)	2234 (86.31±14.98)	_	This study

612 Note: MMU represents minimum mapping units.

614The second highest overlapping rate is approximate 55% in area with Chen's data in 2008615and 2017 (Chen et al., 2021) (Chen et al., 2021). However, the overlapping rate in number is616nearly 45% due to their larger minimum mapping unit (9 pixels). Similarly, a minimum617mapping unit of 55 pixels (50000 m²) in Shugar et al.'s, dataset (2020) (2020) led to the

618 lowest overlap with less than 24% in area. <u>The dataset from Zhang et al. (2015)</u>

619 (2015) ... Zhang's dataset shows fewer glacial lakes in 1990 and 2000 even with a smaller

⁶¹³

620 minimum mapping unit of 3 pixels (Zhang et al., 2015). By inspecting their dataset, we 621 attributed this anomalous discrepancy to a range of glacial lakes that were missinged due to 622 lack of thorough cross-check quality assurance and the limit of a 10-km buffer zone from glaciers during their manual delineation as a result of insufficient high quality images in the 623 624 earlier Landsat era. Our Landsat derived glacial lake dataset has been visually cross-checked 625 over three time periods after the step of object-based automated lake mapping, and also been visually validated by Sentinel-2 derived glacial lakes. Through this series of quality assurance, 626 627 we aim at delivering one of the most reliable multi-decadal glacial lake products for this 628 study area.

629 Other factors, such as minimum mapping units, definition of glacial lakes and study areas, 630 image quality and acquisition dates, mapping methods and quality assurance workflow, might 631 also lead to the discrepancies between the glacial lake datasets. Despite such discrepancies, 632 an increasing number of publically-shared datasets benefit potential users to select the most 633 suitable one for their objectives. Herein, we provide an up-to-date glacial lake dataset derived 634 from both Landsat and Sentinel observations, which further increased the 635 availability of glacial lake datasets promoted the approximation of the size of the sentence of the size of the sentence of the senten

availabilityoptionality of glacial lake datasets promoted the capacity offor GLOFs risk
assessment, and predicting glacier evolutions (Carrivick et al., 2020) cryosphere-hydrological
changes in the context of climate change.

638 6.3 Limitation and updating plan

639 We would like to acknowledge several limitations of our glacier lake dataset, largely due the availability of high quality satellite images in the study area and inadequate field survey data 640 641 (Wang et al., 2020; Chen et al., 2021) (Wang et al., 2020; Chen et al., 2021). First, it is 642 unlikely to collect enough good-quality images within one calendar year for the entire study 643 area due to high possibility of cloud or snow covers. Even though the capacity of repeat 644 observations Even though an capacity of repetitive observations for Landsat-8 OLI and 645 Sentinel-2 increased (Roy et al., 2014; Williamson et al., 2018; Wulder et al., 2019; Paul et al., 2020) (Williamson et al., 2018; Paul et al., 2020; Roy et al., 2014; Wulder et al., 2019), 646 the 2020 glacial lake dataset has to employ images acquired in other years besides 2020. 647 648 Most images used from Landsat and Sentinel platforms were imaged in autumn, and some 649 images taken between April and July and in November also were employed. Distribution and changes in glacial lakes primarily represent the characteristics between August and October. 650 Glacial lakes evolve with time and space (Nie et al., 2017) (Nie et al., 2017), and subtle inter-651 and intra-annual changes (Liu et al., 2020) (Liu et al., 2020) for each time periodin glacial 652 653 lake dataset of each time period were ignored. Second, field investigation data are limited due 654 to low accessibility of high mountain environment in the study area, which restrained the accuracy in classifying the glacial lake types. Although very high-resolution Google Earth 655 656 images were utilized to assist in lake type interpretation, occasional misclassification was 657 inevitable. We implemented two types of classification systems based on a careful utilization of glacier data, DEM, geomorphological features and expert knowledge. However, the lack of 658 659 in situ survey prohibited a thorough validation of the glacial lake types.

7 Data availability

Our glacial lake dataset extracted from Sentinel-2 images in 2020 and Landsat observation between 1990 and 2020 are available online via the Mountain Science Data Center, the Institute of Mountain Hazards and Environment, the Chinese Academy of Sciences at https://doi.org/10.12380/Glaci.msdc.000001 (Lesi et al., 2022)-(Lesi et al., 2022). The glacial lake dataset is provided in both ESRI shapefile format (total size of 22.6 MB) and the Geopackage format (version 1.2.1) with a total size of 9.2MB, which can be opened and further processed by open-source geographic information system software such as QGIS. The glacial lake dataset will be updated using newly collected Landsat and Sentinel images at a five-year interval or modified according to user feedbacks. The updated glacial lake dataset will continue to be released freely and publicly on the Mountain Science Data Center sharing platform.

2 8 Conclusions

Glacial lake inventories of the entire China-Pakistan Economic Corridor in 2020 were
completed based on Landsat and Sentinel-2 images using a human-interactive and
semi-automated mapping method. Both Landsat and Sentinel derived glacial lake datasets
show similar characteristics in spatial distribution and in the statistics of count and area. By
contrast, glacial lake dataset derived from Sentinel-2 images with a spatial resolution of 10 m
has a lower mapping error and more accurate lake boundary than those from 30 m spatial
resolution Landsat images whereas Landsat imagery is more suitable to analyze
spatial-temporal changes at a longer time scalespatial-temporal changes at longer time scale
due to its long-term archived observationsobservation at a consistent spatial resolution of 30
m startingstarted from around 1990.

Glacial lakes in the study area remain relatively stable with a slight increase in number and area between 1990 and 2020 according to Landsat observations. Our dataset reveals that 2154 glacial lakes in 1990 covering 85.1 ± 14.66 km² increased to 2234 lakes with a total area of 86.31 ± 14.98 km². The same mapping method and rigorous workflow of quality assurance and quality control used in this study reduced the error in multi-temporal changes of glacial lakes.

The Hanshaw's error estimation method for automated lake mapping was improved by removing repeatedly calculated edge pixels that vary with lake shape. Therefore, the newly proposed method reduces the estimated value of uncertainty from satellite observations.

692Our glacial lake dataset contains a range of critical parameters that maximize their693potential utility for GLOFs risk evaluation, cryosphere-hydrological and glacier-lake694evolution projection. The dual classification systems of glacial lake types were developed and695are very likely to attract broader researchers and scientists to use our datasets. In comparison696with other existing glacial lake datasets, our products were created through a thorough697consideration of lake types, cross checks and rigorous quality assurance, and will be updated698and released continuously in the data center of mountain science. As such, we expect that our699glacial lake dataset will have significant value to values for cryospheric-hydrology research,700the assessment of glacier-related hazards and engineering project construction in the CPEC.

701 702 703	Supplement. The supplement related to this article is available online.
704 705 706 707 708	Author contributions. ML and YN conceived the study, ML, YN and XD performed data processing and analysis of the glacial lake inventory data, JW contributed to tool development and mapping methods, ML and YN wrote the manuscript. All authors reviewed and edited the manuscript before submission.
709	Competing interests. The authors declare no conflict of interest.
 710 711 712 713 714 715 716 717 718 719 	Acknowledgements. This study was supported by the National Natural Science Foundation of China (Grant Nos. 42171086, 41971153), the International Science & Technology Cooperation Program of China (No. 2018YFE0100100), the Chinese Academy of Sciences "Light of West China" and Natural Sciences and Engineering Research Council of Canada (Grant No. DG-2020-04207).
720	<u>References</u>
721	Ashraf, A., Naz, R., Iqbal, M.B.: Altitudinal dynamics of glacial lakes under changing climate in the Hindu
722	Kush, Karakoram, and Himalaya ranges. Geomorphology, 283: 72-79,
723	https://doi.org/10.1016/j.geomorph.2017.01.033, 2017.
724	Azam, M.F., Kargel, J.S., Shea, J.M., Nepal, S., Haritashya, U.K., Srivastava, S., Maussion, F., Qazi, N.,
725	Chevallier, P., Dimri, A.P., Kulkarni, A.V., Cogley, J.G., Bahuguna, I.: Glaciohydrology of the
726	Himalaya-Karakoram. Science, 373: eabf3668, https://doi.org/10.1126/science.abf3668, 2021.
727	Battamo, A.Y., Varis, O., Sun, P., Yang, Y., Oba, B.T., Zhao, L.: Mapping socio-ecological resilience along the
728	seven economic corridors of the Belt and Road Initiative. J. Clean. Prod., 309: 127341,
729	https://doi.org/10.1016/j.jclepro.2021.127341, 2021.
730	Bhambri, R., Hewitt, K., Kawishwar, P., Kumar, A., Verma, A., Snehmani, Tiwari, S., Misra, A.: Ice-dams,
731	outburst floods, and movement heterogeneity of glaciers, Karakoram. Global Planet. Change, 180: 100-116,
732	https://doi.org/10.1016/j.gloplacha.2019.05.004, 2019.

733	Bhattacharya, A., Bolch, T., Mukherjee, K., King, O., Menounos, B., Kapitsa, V., Neckel, N., Yang, W., Yao, T.:
734	High Mountain Asian glacier response to climate revealed by multi-temporal satellite observations since the
735	<u>1960s. Nat. Commun., 12: 4133, https://doi.org/10.1038/s41467-021-24180-y, 2021.</u>
736	Bolch, T., Pieczonka, T., Mukherjee, K., Shea, J.: Brief communication: Glaciers in the Hunza catchment
737	(Karakoram) have been nearly in balance since the 1970s. The Cryosphere, 11: 531-539,
738	https://doi.org/10.5194/tc-11-531-2017, 2017.
739	Brun, F., Berthier, E., Wagnon, P., K ääb, A., Treichler, D.: A spatially resolved estimate of High Mountain Asia
740	glacier mass balances from 2000 to 2016. Nat. Geosci., 10: 668-673, https://doi.org/10.1038/ngeo2999, 2017.
741	Brun, F., Wagnon, P., Berthier, E., Jomelli, V., Maharjan, S.B., Shrestha, F., Kraaijenbrink, P.D.A.:
742	Heterogeneous Influence of Glacier Morphology on the Mass Balance Variability in High Mountain Asia. J.
743	Geophys. Res-Earth, 124: 1331-1345, https://doi.org/10.1029/2018JF004838, 2019.
744	Carrivick, J.L., Tweed, F.S.: Proglacial lakes: character, behaviour and geological importance. Quaternary Sci.
745	Rev., 78: 34-52, https://doi.org/10.1016/j.quascirev.2013.07.028, 2013.
746	Carrivick, J.L., Quincey, D.J.: Progressive increase in number and volume of ice-marginal lakes on the western
747	margin of the Greenland Ice Sheet. Global Planet. Change, 116: 156-163,
748	https://doi.org/10.1016/j.gloplacha.2014.02.009, 2014.
749	Carrivick, J.L., Tweed, F.S.: A global assessment of the societal impacts of glacier outburst floods. Global
750	Planet. Change, 144: 1-16, https://doi.org/10.1016/j.gloplacha.2016.07.001, 2016.
751	Carrivick, J.L., Tweed, F.S., Sutherland, J.L., Mallalieu, J.: Toward Numerical Modeling of Interactions
752	Between Ice-Marginal Proglacial Lakes and Glaciers. Front. Earth Sci. 8.
753	https://doi.org/10.3389/feart.2020.577068, 2020.
754	Chen, F., Zhang, M., Guo, H., Allen, S., Kargel, J.S., Haritashya, U.K., Watson, C.S.: Annual 30 m dataset for

755	glacial lakes in High Mountain Asia from 2008 to 2017. Earth System Science Data, 13: 741-766,
756	https://doi.org/10.5194/essd-13-741-2021, 2021.
757	Chen, X., Cui, P., You, Y., Cheng, Z., Khan, A., Ye, C., Zhang, S.: Dam-break risk analysis of the Attabad
758	landslide dam in Pakistan and emergency countermeasures. Landslides, 14: 675-683,
759	https://doi.org/10.1007/s10346-016-0721-7, 2017.
760	Emmer, A., Cuřín, V.: Can a dam type of an alpine lake be derived from lake geometry? A negative result. J. Mt.
761	SciEngl., 18: 614-621, https://doi.org/10.1007/s11629-020-6003-9, 2021.
762	Farr, T.G., Rosen, P.A., Caro, E., Crippen, R., Duren, R., Hensley, S., Kobrick, M., Paller, M., Rodriguez, E.,
763	Roth, L., Seal, D., Shaffer, S., Shimada, J., Umland, J., Werner, M., Oskin, M., Burbank, D., Alsdorf, D.: The
764	Shuttle Radar Topography Mission. Rev. Geophys., 45: RG2004, https://doi.org/10.1029/2005RG000183, 2007.
765	Gardelle, J., Arnaud, Y., Berthier, E.: Contrasted evolution of glacial lakes along the Hindu Kush Himalaya
766	mountain range between 1990 and 2009. Global Planet. Change, 75: 47-55,
767	https://doi.org/10.1016/j.gloplacha.2010.10.003, 2011.
767 768	https://doi.org/10.1016/j.gloplacha.2010.10.003, 2011. Hanshaw, M.N., Bookhagen, B.: Glacial areas, lake areas, and snow lines from 1975 to 2012: status of the
768	Hanshaw, M.N., Bookhagen, B.: Glacial areas, lake areas, and snow lines from 1975 to 2012: status of the
768 769	Hanshaw, M.N., Bookhagen, B.: Glacial areas, lake areas, and snow lines from 1975 to 2012: status of the Cordillera Vilcanota, including the Quelccaya Ice Cap, northern central Andes, Peru. The Cryosphere, 8:
768 769 770	Hanshaw, M.N., Bookhagen, B.: Glacial areas, lake areas, and snow lines from 1975 to 2012: status of the Cordillera Vilcanota, including the Quelccaya Ice Cap, northern central Andes, Peru. The Cryosphere, 8: 359-376, https://doi.org/10.5194/tc-8-359-2014, 2014.
768 769 770 771	Hanshaw, M.N., Bookhagen, B.: Glacial areas, lake areas, and snow lines from 1975 to 2012: status of the Cordillera Vilcanota, including the Quelccaya Ice Cap, northern central Andes, Peru. The Cryosphere, 8: 359-376, https://doi.org/10.5194/tc-8-359-2014, 2014, Hewitt, K.: The Karakoram Anomaly? Glacier Expansion and the 'Elevation Effect,' Karakoram Himalaya. Mt.
768 769 770 771 772	Hanshaw, M.N., Bookhagen, B.: Glacial areas, lake areas, and snow lines from 1975 to 2012: status of the Cordillera Vilcanota, including the Quelccaya Ice Cap, northern central Andes, Peru. The Cryosphere, 8: 359-376, https://doi.org/10.5194/tc-8-359-2014, 2014, Hewitt, K.: The Karakoram Anomaly? Glacier Expansion and the 'Elevation Effect,' Karakoram Himalaya. Mt. Res. Dev., 25: 332-340, https://doi.org/10.1659/0276-4741(2005)025[0332:TKAGEA]2.0.CO;2, 2005.
 768 769 770 771 772 773 	Hanshaw, M.N., Bookhagen, B.: Glacial areas, lake areas, and snow lines from 1975 to 2012: status of the. Cordillera Vilcanota, including the Quelccaya Ice Cap, northern central Andes, Peru. The Cryosphere, 8: 359-376, https://doi.org/10.5194/tc-8-359-2014, 2014. Hewitt, K.: The Karakoram Anomaly? Glacier Expansion and the 'Elevation Effect,' Karakoram Himalaya. Mt. Res. Dev., 25: 332-340, https://doi.org/10.1659/0276-4741(2005)025[0332:TKAGEA]2.0.CO;2, 2005. Hewitt, K., 2014. Glaciers of the Karakoram Himalaya: Glacial Environments, Processes, Hazards and

777	RepUK, 11: 4481, https://doi.org/10.1038/s41598-021-83509-1, 2021.
778	Huggel, C., K ääb, A., Haeberli, W., Teysseire, P., Paul, F.: Remote sensing based assessment of hazards from
779	glacier lake outbursts: a case study in the Swiss Alps. Can. Geotech. J., 39: 316-330,
780	https://doi.org/10.1139/t01-099, 2002.
781	Hugonnet, R., Mcnabb, R., Berthier, E., Menounos, B., Nuth, C., Girod, L., Farinotti, D., Huss, M., Dussaillant,
782	I., Brun, F., K ääb, A.: Accelerated global glacier mass loss in the early twenty-first century. Nature, 592:
783	726-731, https://doi.org/10.1038/s41586-021-03436-z, 2021.
784	Jarvis, A., Reuter, H.I., Nelson, A., Guevara, E., 2008. Hole-filled seamless SRTM data V4. 2008, International
785	Centre for Tropical Agriculture (CIAT), available from http://srtm.csi.cgiar.org.
786	Jiang, S., Nie, Y., Liu, Q., Wang, J., Liu, L., Hassan, J., Liu, X., Xu, X.: Glacier Change, Supraglacial Debris
787	Expansion and Glacial Lake Evolution in the Gyirong River Basin, Central Himalayas, between 1988 and 2015.
-	
788	Remote SensBasel, 10: 986, https://doi.org/10.3390/rs10070986, 2018.
788	<u>Remote SensBasel, 10: 986, https://doi.org/10.3390/rs10070986, 2018.</u> <u>K ääb, A., Berthier, E., Nuth, C., Gardelle, J., Arnaud, Y.: Contrasting patterns of early twenty-first-century</u>
789	Kääb, A., Berthier, E., Nuth, C., Gardelle, J., Arnaud, Y.: Contrasting patterns of early twenty-first-century
789 790	K ääb, A., Berthier, E., Nuth, C., Gardelle, J., Arnaud, Y.: Contrasting patterns of early twenty-first-century_ glacier mass change in the Himalayas. Nature, 488: 495-498, https://doi.org/10.1038/nature11324, 2012.
789 790 791	 <u>K ääb, A., Berthier, E., Nuth, C., Gardelle, J., Arnaud, Y.: Contrasting patterns of early twenty-first-century</u> <u>glacier mass change in the Himalayas. Nature, 488: 495-498, https://doi.org/10.1038/nature11324, 2012.</u> <u>Lesi, M., Nie, Y., Shugar, D.H., Wang, J., Deng, Q., Chen, H.: Landsat and Sentinel-derived glacial lake dataset</u>
789 790 791 792	 <u>K ääb, A., Berthier, E., Nuth, C., Gardelle, J., Arnaud, Y.: Contrasting patterns of early twenty-first-century</u> <u>glacier mass change in the Himalayas. Nature, 488: 495-498, https://doi.org/10.1038/nature11324, 2012.</u> <u>Lesi, M., Nie, Y., Shugar, D.H., Wang, J., Deng, Q., Chen, H.: Landsat and Sentinel-derived glacial lake dataset</u> <u>in the China-Pakistan Economic Corridor from 1990 to 2020. Mountain Science Data Center,</u>
 789 790 791 792 793 	K ääb, A., Berthier, E., Nuth, C., Gardelle, J., Arnaud, Y.: Contrasting patterns of early twenty-first-century glacier mass change in the Himalayas. Nature, 488: 495-498, https://doi.org/10.1038/nature11324, 2012. Lesi, M., Nie, Y., Shugar, D.H., Wang, J., Deng, Q., Chen, H.: Landsat and Sentinel-derived glacial lake dataset in the China-Pakistan Economic Corridor from 1990 to 2020. Mountain Science Data Center, https://doi.org/10.12380/Glaci.msdc.000001 CSTR:1a006.11.Glaci.msdc.000001, 2022.
 789 790 791 792 793 794 	K ääb, A., Berthier, E., Nuth, C., Gardelle, J., Arnaud, Y.: Contrasting patterns of early twenty-first-century glacier mass change in the Himalayas. Nature, 488: 495-498, https://doi.org/10.1038/nature11324, 2012. Lesi, M., Nie, Y., Shugar, D.H., Wang, J., Deng, Q., Chen, H.: Landsat and Sentinel-derived glacial lake dataset in the China-Pakistan Economic Corridor from 1990 to 2020. Mountain Science Data Center, https://doi.org/10.12380/Glaci.msdc.000001 CSTR:1a006.11.Glaci.msdc.000001, 2022. Li, D., Shangguan, D., Anjum, M.N.: Glacial Lake Inventory Derived from Landsat 8 OLI in 2016–2018 in
 789 790 791 792 793 794 795 	K ääb, A., Berthier, E., Nuth, C., Gardelle, J., Arnaud, Y.: Contrasting patterns of early twenty-first-century. glacier mass change in the Himalayas. Nature, 488: 495-498, https://doi.org/10.1038/nature11324, 2012. Lesi, M., Nie, Y., Shugar, D.H., Wang, J., Deng, Q., Chen, H.: Landsat and Sentinel-derived glacial lake dataset in the China-Pakistan Economic Corridor from 1990 to 2020. Mountain Science Data Center, https://doi.org/10.12380/Glaci.msdc.000001CSTR:1a006.11.Glaci.msdc.000001, 2022. Li, D., Shangguan, D., Anjum, M.N.: Glacial Lake Inventory Derived from Landsat 8 OLI in 2016–2018 in China–Pakistan Economic Corridor. ISPRS international journal of geo-information, 9: 294,

799	https://doi.org/10.1016/j.jclepro.2020.125406, 2021.
800	Liu, Q., Mayer, C.: Distribution and interannual variability of supraglacial lakes on debris-covered glaciers in
801	the Khan Tengri-Tumor Mountains, Central Asia. Environ. Res. Lett., 10: 014014 2015.
802	Liu, Q., Mayer, C., Wang, X., Nie, Y., Wu, K., Wei, J., Liu, S.: Interannual flow dynamics driven by frontal
803	retreat of a lake-terminating glacier in the Chinese Central Himalaya. Earth Planet. Sc. Lett., 546: 116450,
804	https://doi.org/10.1016/j.epsl.2020.116450, 2020.
805	Lutz, A.F., Immerzeel, W.W., Shrestha, A.B., Bierkens, M.F.P.: Consistent increase in High Asia's runoff due to
806	increasing glacier melt and precipitation. Nat. Clim. Change, 4: 587-592, https://doi.org/10.1038/nclimate2237,
807	<u>2014.</u>
808	Lyons, E.A., Sheng, Y., Smith, L.C., Li, J., Hinkel, K.M., Lenters, J.D., Wang, J.: Quantifying sources of error
809	in multitemporal multisensor lake mapping. Int. J. Remote Sens., 34: 7887-7905,
810	https://doi.org/10.1080/01431161.2013.827343, 2013.
811	Mart ń, C.N.S., Ponce, J.F., Montes, A., Balocchi, L.D., Gorza, C., Andrea, C.: Proglacial landform assemblage
812	in a rapidly retreating cirque glacier due to temperature increase since 1970, Fuegian Andes, Argentina.
813	Geomorphology, 390: 107861, https://doi.org/10.1016/j.geomorph.2021.107861, 2021.
814	Maurer, J.M., Schaefer, J.M., Rupper, S., Corley, A.: Acceleration of ice loss across the Himalayas over the past
815	40 years. Sci. Adv., 5: eaav7266, https://doi.org/10.1126/sciadv.aav7266, 2019.
816	Mcfeeters, S.K.: The use of the Normalized Difference Water Index (NDWI) in the delineation of open water
817	features. Int. J. Remote Sens., 17: 1425 - 1432 1996.
818	Miles, E.S., Watson, C.S., Brun, F., Berthier, E., Esteves, M., Quincey, D.J., Miles, K.E., Hubbard, B., Wagnon,
819	P.: Glacial and geomorphic effects of a supraglacial lake drainage and outburst event, Everest region, Nepal
820	Himalaya. The Cryosphere, 12: 3891-3905, https://doi.org/10.5194/tc-12-3891-2018, 2018.

821	Nie, Y., Zhang, Y., Liu, L., Zhang, J.: Glacial change in the vicinity of Mt. Qomolangma (Everest), central high
822	Himalayas since 1976. J. Geogr. Sci., 20: 667-686, https://doi.org/10.1007/s11442-010-0803-8, 2010.
823	Nie, Y., Sheng, Y., Liu, Q., Liu, L., Liu, S., Zhang, Y., Song, C.: A regional-scale assessment of Himalayan
824	glacial lake changes using satellite observations from 1990 to 2015. Remote Sens. Environ., 189: 1-13,
825	https://doi.org/10.1016/j.rse.2016.11.008, 2017.
826	Nie, Y., Liu, Q., Wang, J., Zhang, Y., Sheng, Y., Liu, S.: An inventory of historical glacial lake outburst floods
827	in the Himalayas based on remote sensing observations and geomorphological analysis. Geomorphology, 308:
828	91-106, https://doi.org/10.1016/j.geomorph.2018.02.002, 2018.
829	Nie, Y., Liu, W., Liu, Q., Hu, X., Westoby, M.J.: Reconstructing the Chongbaxia Tsho glacial lake outburst
830	flood in the Eastern Himalaya: Evolution, process and impacts. Geomorphology, 370: 107393,
831	https://doi.org/10.1016/j.geomorph.2020.107393, 2020.
832	Nie, Y., Pritchard, H.D., Liu, Q., Hennig, T., Wang, W., Wang, X., Liu, S., Nepal, S., Samyn, D., Hewitt, K.,
833	Chen, X.: Glacial change and hydrological implications in the Himalaya and Karakoram. Nature Reviews Earth
834	& Environment, 2: 91-106, https://doi.org/10.1038/s43017-020-00124-w, 2021.
835	Paul, F., Rastner, P., Azzoni, R.S., Diolaiuti, G., Fugazza, D., Le Bris, R., Nemec, J., Rabatel, A., Ramusovic,
836	M., Schwaizer, G., Smiraglia, C.: Glacier shrinkage in the Alps continues unabated as revealed by a new glacier
837	inventory from Sentinel-2. Earth System Science Data, 12: 1805-1821,
838	https://doi.org/10.5194/essd-12-1805-2020, 2020.
839	Pfeffer, W.T., Arendt, A.A., Bliss, A., Bolch, T., Cogley, J.G., Gardner, A.S., Hagen, J., Hock, R., Kaser, G.,
840	Kienholz, C., Miles, E.S., Moholdt, G., Mölg, N., Paul, F., Radić, V., Rastner, P., Raup, B.H., Rich, J., Sharp,
841	M.J.: The Randolph Glacier Inventory: a globally complete inventory of glaciers. J. Glaciol., 60: 537-552,
842	https://doi.org/10.3189/2014JoG13J176, 2014.

843	Post, A., Mayo, L.R., 1971. Glacier dammed lakes and outburst floods in Alaska: U.S. Geological Survey
844	Hydrologic Investigations Atlas 455, U.S. Geological Survey.
845	Pritchard, H.D.: Asia's shrinking glaciers protect large populations from drought stress. Nature, 569: 649-654,
846	https://doi.org/10.1038/s41586-019-1240-1, 2019.
847	Quincey, D.J., Richardson, S.D., Luckman, A., Lucas, R.M., Reynolds, J.M., Hambrey, M.J., Glasser, N.F.:
848	Early recognition of glacial lake hazards in the Himalaya using remote sensing datasets. Global Planet. Change,
849	56: 137-152, https://doi.org/10.1016/j.gloplacha.2006.07.013, 2007.
850	Rabus, B., Eineder, M., Roth, A., Bamler, R.: The shuttle radar topography mission-a new class of digital
851	elevation models acquired by spaceborne radar. ISPRS J. Photogramm., 57: 241-262,
852	https://doi.org/10.1016/S0924-2716(02)00124-7, 2003.
853	RGI Consortium: Randolph Glacier Inventory – A Dataset of Global Glacier Outlines: Version 6.0: Technical
854	Report, https://doi.org/10.7265/N5-RGI-60, 2017.
855	Rick, B., Mcgrath, D., Armstrong, W., Mccoy, S.W.: Dam type and lake location characterize ice-marginal lake
856	area change in Alaska and NW Canada between 1984 and 2019. The Cryosphere, 16: 297-314,
857	https://doi.org/10.5194/tc-16-297-2022, 2022.
858	Rounce, D.R., Hock, R., Shean, D.E.: Glacier Mass Change in High Mountain Asia Through 2100 Using the
859	Open-Source Python Glacier Evolution Model (PyGEM). Front. Earth Sci. 7: 331.
860	https://doi.org/10.3389/feart.2019.00331, 2020.
861	Roy, D.P., Wulder, M.A., Loveland, T.R., C. E., W., Allen, R.G., Anderson, M.C., Helder, D., Irons, J.R.,
862	Johnson, D.M., Kennedy, R., Scambos, T.A., Schaaf, C.B., Schott, J.R., Sheng, Y., Vermote, E.F., Belward,
863	A.S., Bindschadler, R., Cohen, W.B., Gao, F., Hipple, J.D., Hostert, P., Huntington, J., Justice, C.O., Kilic, A.,
864	Kovalskyy, V., Lee, Z.P., Lymburner, L., Masek, J.G., Mccorkel, J., Shuai, Y., Trezza, R., Vogelmann, J.,

865	Wynne, R.H., Zhu, Z.: Landsat-8: Science and product vision for terrestrial global change research. Remote
866	Sens. Environ., 145: 154-172, https://doi.org/10.1016/j.rse.2014.02.001, 2014.
867	Sakai, A.: Brief communication: Updated GAMDAM glacier inventory over high-mountain Asia. The
868	Cryosphere, 13: 2043-2049, https://doi.org/10.5194/tc-13-2043-2019, 2019.
869	Shean, D.E., Bhushan, S., Montesano, P., Rounce, D.R., Arendt, A., Osmanoglu, B.: A Systematic, Regional
870	Assessment of High Mountain Asia Glacier Mass Balance. Front. Earth Sci, 7: 363,
871	https://doi.org/10.3389/feart.2019.00363, 2020.
872	Sheng, Y., Song, C., Wang, J., Lyons, E.A., Knox, B.R., Cox, J.S., Gao, F.: Representative lake water extent
873	mapping at continental scales using multi-temporal Landsat-8 imagery. Remote Sens. Environ., 185: 129-141,
874	https://doi.org/10.1016/j.rse.2015.12.041, 2016.
875	Shugar, D.H., Burr, A., Haritashya, U.K., Kargel, J.S., Watson, C.S., Kennedy, M.C., Bevington, A.R., Betts,
876	R.A., Harrison, S., Strattman, K.: Rapid worldwide growth of glacial lakes since 1990. Nat. Clim. Change, 10:
877	939-945, https://doi.org/10.1038/s41558-020-0855-4, 2020.
878	Shugar, D.H., Jacquemart, M., Shean, D., Bhushan, S., Upadhyay, K., Sattar, A., Schwanghart, W., Mcbride, S.,
879	de Vries, M., Mergili, M., Emmer, A., Deschamps-Berger, C., Mcdonnell, M., Bhambri, R., Allen, S., Berthier,
880	E., Carrivick, J.L., Clague, J.J., Dokukin, M., Dunning, S.A., Frey, H., Gascoin, S., Haritashya, U.K., Huggel,
881	C., Kaab, A., Kargel, J.S., Kavanaugh, J.L., Lacroix, P., Petley, D., Rupper, S., Azam, M.F., Cook, S.J., Dimri,
882	A.P., Eriksson, M., Farinotti, D., Fiddes, J., Gnyawali, K.R., Harrison, S., Jha, M., Koppes, M., Kumar, A.,
883	Leinss, S., Majeed, U., Mal, S., Muhuri, A., Noetzli, J., Paul, F., Rashid, I., Sain, K., Steiner, J., Ugalde, F.,
884	Watson, C.S., Westoby, M.J.: A massive rock and ice avalanche caused the 2021 disaster at Chamoli, Indian
885	Himalaya. Science, 373: 300-306, https://doi.org/10.1126/science.abh4455, 2021.
886	Ullah, S., You, Q., Ali, A., Ullah, W., Jan, M.A., Zhang, Y., Xie, W., Xie, X.: Observed changes in maximum

887	and minimum temperatures over China- Pakistan economic corridor during 1980–2016. Atmos. Res., 216:
888	37-51, https://doi.org/10.1016/j.atmosres.2018.09.020, 2019.
889	Viviroli, D., Kummu, M., Meybeck, M., Kallio, M., Wada, Y.: Increasing dependence of lowland populations
890	on mountain water resources. Nature Sustainability, 3: 917-928, https://doi.org/10.1038/s41893-020-0559-9,
891	<u>2020.</u>
892	Wang, J., Sheng, Y., Tong, T.S.D.: Monitoring decadal lake dynamics across the Yangtze Basin downstream of
893	Three Gorges Dam. Remote Sens. Environ., 152: 251-269, https://doi.org/10.1016/j.rse.2014.06.004, 2014.
894	Wang, J., Sheng, Y., Wada, Y.: Little impact of the Three Gorges Dam on recent decadal lake decline across
895	China's Yangtze Plain. Water Resour. Res., 53: 3854-3877, https://doi.org/10.1002/2016WR019817, 2017.
896	Wang, J., Song, C., Reager, J.T., Yao, F., Famiglietti, J.S., Sheng, Y., Macdonald, G.M., Brun, F., Schmied,
897	H.M., Marston, R.A., Wada, Y.: Recent global decline in endorheic basin water storages. Nat. Geosci., 11:
898	926-932, https://doi.org/10.1038/s41561-018-0265-7, 2018.
898 899	926-932, https://doi.org/10.1038/s41561-018-0265-7, 2018. Wang, X., Ding, Y., Liu, S., Jiang, L., Wu, K., Jiang, Z., Guo, W.: Changes of glacial lakes and implications in
899	Wang, X., Ding, Y., Liu, S., Jiang, L., Wu, K., Jiang, Z., Guo, W.: Changes of glacial lakes and implications in
899 900	Wang, X., Ding, Y., Liu, S., Jiang, L., Wu, K., Jiang, Z., Guo, W.: Changes of glacial lakes and implications in <u>Tian Shan, Central Asia, based on remote sensing data from 1990 to 2010. Environ. Res. Lett., 8: 44052,</u>
899 900 901	Wang, X., Ding, Y., Liu, S., Jiang, L., Wu, K., Jiang, Z., Guo, W.: Changes of glacial lakes and implications in Tian Shan, Central Asia, based on remote sensing data from 1990 to 2010. Environ. Res. Lett., 8: 44052, https://doi.org/10.1088/1748-9326/8/4/044052, 2013.
899900901902	 Wang, X., Ding, Y., Liu, S., Jiang, L., Wu, K., Jiang, Z., Guo, W.: Changes of glacial lakes and implications in Tian Shan, Central Asia, based on remote sensing data from 1990 to 2010. Environ. Res. Lett., 8: 44052, https://doi.org/10.1088/1748-9326/8/4/044052, 2013. Wang, X., Liu, S., Zhang, J.: A new look at roles of the cryosphere in sustainable development. Advances in
899900901902903	 Wang, X., Ding, Y., Liu, S., Jiang, L., Wu, K., Jiang, Z., Guo, W.: Changes of glacial lakes and implications in <u>Tian Shan, Central Asia, based on remote sensing data from 1990 to 2010. Environ. Res. Lett., 8: 44052,</u> <u>https://doi.org/10.1088/1748-9326/8/4/044052, 2013.</u> Wang, X., Liu, S., Zhang, J.: A new look at roles of the cryosphere in sustainable development. Advances in <u>Climate Change Research, 10: 124-131, https://doi.org/10.1016/j.accre.2019.06.005, 2019.</u>
 899 900 901 902 903 904 	 Wang, X., Ding, Y., Liu, S., Jiang, L., Wu, K., Jiang, Z., Guo, W.: Changes of glacial lakes and implications in Tian Shan, Central Asia, based on remote sensing data from 1990 to 2010. Environ. Res. Lett., 8: 44052, https://doi.org/10.1088/1748-9326/8/4/044052, 2013. Wang, X., Liu, S., Zhang, J.: A new look at roles of the cryosphere in sustainable development. Advances in Climate Change Research, 10: 124-131, https://doi.org/10.1016/j.accre.2019.06.005, 2019. Wang, X., Guo, X., Yang, C., Liu, Q., Wei, J., Zhang, Y., Liu, S., Zhang, Y., Jiang, Z., Tang, Z.: Glacial lake.
 899 900 901 902 903 904 905 	 Wang, X., Ding, Y., Liu, S., Jiang, L., Wu, K., Jiang, Z., Guo, W.: Changes of glacial lakes and implications in Tian Shan, Central Asia, based on remote sensing data from 1990 to 2010. Environ. Res. Lett., 8: 44052, https://doi.org/10.1088/1748-9326/8/4/044052, 2013. Wang, X., Liu, S., Zhang, J.: A new look at roles of the cryosphere in sustainable development. Advances in Climate Change Research, 10: 124-131, https://doi.org/10.1016/j.accre.2019.06.005, 2019. Wang, X., Guo, X., Yang, C., Liu, Q., Wei, J., Zhang, Y., Liu, S., Zhang, Y., Jiang, Z., Tang, Z.: Glacial lake inventory of high-mountain Asia in 1990 and 2018 derived from Landsat images. Earth System Science Data,

909 https://doi.org/https://doi.org/10.1016/j.srs.2020.100008, 2020.
--

- 910 Westoby, M.J., Glasser, N.F., Brasington, J., Hambrey, M.J., Quincey, D.J., Reynolds, J.M.: Modelling outburst
- 911 floods from moraine-dammed glacial lakes. Earth-Sci. Rev., 134: 137-159,
- 912 <u>https://doi.org/10.1016/j.earscirev.2014.03.009, 2014.</u>

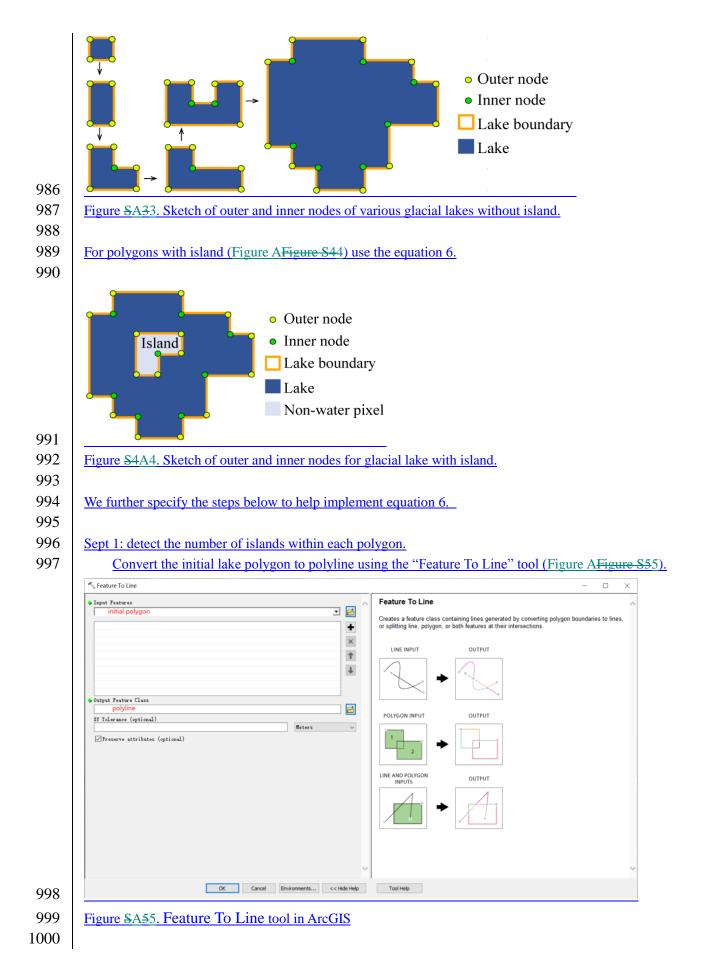
- 913 Williamson, A.G., Banwell, A.F., Willis, I.C., Arnold, N.S.: Dual-satellite (Sentinel-2 and Landsat 8) remote
- 914 <u>sensing of supraglacial lakes in Greenland. The Cryosphere, 12: 3045-3065,</u>
- 915 <u>https://doi.org/10.5194/tc-12-3045-2018, 2018.</u>
- 916 Wulder, M.A., Loveland, T.R., Roy, D.P., Crawford, C.J., Masek, J.G., Woodcock, C.E., Allen, R.G., Anderson,
- 917 M.C., Belward, A.S., Cohen, W.B., Dwyer, J., Erb, A., Gao, F., Griffiths, P., Helder, D., Hermosilla, T., Hipple,
- 918 J.D., Hostert, P., Hughes, M.J., Huntington, J., Johnson, D.M., Kennedy, R., Kilic, A., Li, Z., Lymburner, L.,
- 919 Mccorkel, J., Pahlevan, N., Scambos, T.A., Schaaf, C., Schott, J.R., Sheng, Y., Storey, J., Vermote, E.,
- 920 Vogelmann, J., White, J.C., Wynne, R.H., Zhu, Z.: Current status of Landsat program, science, and applications.
- 921 <u>Remote Sens. Environ., 225: 127-147, https://doi.org/https://doi.org/10.1016/j.rse.2019.02.015, 2019.</u>
- 922 Yao, C., Wang, X., Zhao, X., Wei, J., Zhang, Y.: Temporal and Spatial Changes of Glacial Lakes in the
- 923 China-Pakistan Economic Corridor from 1990 to 2018. Journal of Glaciology and Geocryology, 42: 33-42,
- 924 <u>https://doi.org/https://doi.org/10.7522/j.issn.1000-0240.2020.0009, 2020.</u>
- 925 Yao, T., Thompson, L., Yang, W., Yu, W.S., Gao, Y., Guo, X.J., Yang, X.X., Duan, K.Q., Zhao, H.B., Xu, B.Q.,
- 926 Pu, J.C., Lu, A.X., Xiang, Y., Kattel, D.B., Joswiak, D.: Different glacier status with atmospheric circulations in
- 927 <u>Tibetan Plateau and surroundings. Nat. Clim. Change, 2: 663-667, https://doi.org/10.1038/NCLIMATE1580,</u>
- 928 <u>2012.</u>
- 929 Yao, X., Liu, S., Han, L., Sun, M., Zhao, L.: Definition and classification system of glacial lake for inventory
- 930 and hazards study. J. Geogr. Sci., 28: 193-205, https://doi.org/10.1007/s11442-018-1467-z, 2018.

931	Zhang, G., Yao, T., Xie, H., Wang, W., Yang, W.: An inventory of glacial lakes in the Third Pole region and
932	their changes in response to global warming. Global Planet. Change, 131: 148-157,
933	https://doi.org/10.1016/j.gloplacha.2015.05.013, 2015.
934	Zhang, M., Chen, F., Tian, B.: An automated method for glacial lake mapping in High Mountain Asia using
935	Landsat 8 imagery. J. Mt. SciEngl., 15: 13-24, https://doi.org/10.1007/s11629-017-4518-5, 2018.
936	Zhao, W., Xiong, D., Wen, F., Wang, X.: Lake area monitoring based on land surface temperature in the Tibetan
937	Plateau from 2000 to 2018. Environ. Res. Lett., 15, https://doi.org/10.1088/1748-9326/ab9b41, 2020.
938	Zheng, G., Allen, S.K., Bao, A., Ballesteros-Cánovas, J.A., Huss, M., Zhang, G., Li, J., Yuan, Y., Jiang, L., Yu,
939	T., Chen, W., Stoffel, M.: Increasing risk of glacial lake outburst floods from future Third Pole deglaciation. Nat.
940	Clim. Change, 11: 411-417, https://doi.org/10.1038/s41558-021-01028-3, 2021.
941 942	

<u>Appendix</u> <u>Tutorial for Improved Uncertai</u>	nty Estimating M	ethod
The Hanshaw's equation was	originally propo	osed for pixelated polygons (such as a polygon direct
		erformed more robustly than manually digitized polygo
		xel edges). Our improved method also performs better f
		o helping implement our improved uncertainty estimation
method.		
Procedure of uncertainty estim	ating method (usi	ng ArcGIS for example)
. Removing redundant nodes		<u> </u>
		pixelated lake polygons (directly extracted from satell
		he value of inner nodes. If no redundant nodes exist, the
		the "Simplify Polygon" tool in ArcGIS to remove the
nodes (Figure A1 Figure S1).		
n the Simplify Polygon panel		
Input your dataset.		
Set the output path and ou	ttput file name.	
		commended "POINT REMOVE".
		ithm. In this step, we need to ensure that the polyg
	•	undant nodes. Generally, a tolerance of 1 meter will suffi
or you can adjust the threshold		·
Simplify Polygon		X
	<u>^</u>	Keep collapsed points (optional)
🖕 Input Features 🤍		keep collapsed points (optional)
	▼ 6	Specifies whether to create an output point feature class to store the centers of polygons that
• Output Feature Class 2		Specifies whether to create an output point feature class to store the centers of polygons that are removed because they are smaller than the Minimum area parameter. The point output is ferived; it will use the same name and location as the Output feature class parameter but
• Output Feature Class ? Simplification Algorithm FOINT_ERMOVE Simplification Tolerance		Specifies whether to create an output point feature class to store the centers of polygons that are removed because they are smaller than the Minimum area parameter. The point output is derived, it will use the same name and location as the Output feature class parameter but with a _Pnt suffix.
Simplification Algorithm FOINT_REMOVE Simplification Tolerance Innimum Area (optional)	Eaters	Specifies whether to create an output point feature class to store the centers of polygons that are removed because they are smaller than the Minimum area parameter. The point output is derived; it will use the same name and location as the Output feature class parameter but with a _Pnt suffix.
Output Feature Class Sisplification Algorithm POINT_REMOVE Simplification Tolerance Intinuum Area (optional) Nandling Topological Errors (optional)	E	Specifies whether to create an output point feature class to store the centers of polygons that are removed because they are smaller than the Minimum area parameter. The point output is derived; it will use the same name and location as the Output feature class parameter but with a _Pnt suffix.
Output Feature Class Simplification Algorithm FOINT_REMOVE Simplification Tolerance Innimum Area (optional) 0	Eaters	Specifies whether to create an output point feature class to store the centers of polygons that are removed because they are smaller than the Minimum area parameter. The point output is derived; it will use the same name and location as the Output feature class parameter but with a _Pnt suffix.
Output Feature Class 2 Simplification Algorithm POINT_REMOVE Simplification Tolerance	Keters Square Meters V	Specifies whether to create an output point feature class to store the centers of polygons that are removed because they are smaller than the Minimum area parameter. The point output is derived; it will use the same name and location as the Output feature class parameter but with a _Pnt suffix.
Output Feature Class Simplification Algorithm FOINT_REMOVE Simplification Tolerance International Internation Internatio Internation Internation Internation Internation Int	E	Specifies whether to create an output point feature class to store the centers of polygons that are removed because they are smaller than the Minimum area parameter. The point output is derived; it will use the same name and location as the Output feature class parameter but with a _Pnt suffix.
Output Feature Class Simplification Algorithm FOINT_REMOVE Simplification Tolerance Infinimum Area (optional) Mundling Topological Errors (optional) RESOLVE_REBORS Reep collapsed points (optional)		Specifies whether to create an output point feature class to store the centers of polygons that are removed because they are smaller than the Minimum area parameter. The point output is derived; it will use the same name and location as the Output feature class parameter but with a _Pnt suffix.
Output Feature Class Sisplification Algorithm FOINT_REMOVE Sisplification Tolerance Iminium Area (optional) Mandling Topological Errors (optional) RESOLVE_ERRORS Keep collapsed points (optional)		Specifies whether to create an output point feature class to store the centers of polygons that are removed because they are smaller than the Minimum area parameter. The point output is derived; it will use the same name and location as the Output feature class parameter but with a _Pnt suffix.
Output Feature Class Sisplification Algorithm FOINT_REMOVE Sisplification Tolerance Iminium Area (optional) Mandling Topological Errors (optional) RESOLVE_ERRORS Keep collapsed points (optional)		Specifies whether to create an output point feature class to store the centers of polygons that are removed because they are smaller than the Minimum area parameter. The point output is derived; it will use the same name and location as the Output feature class parameter but with a _Pnt suffix.
Output Feature Class Simplification Algorithm FOINT_REMOVE Simplification Tolerance		Specifies whether to create an output point feature class to store the centers of polygons that are removed because they are smaller than the Minimum area parameter. The point output is derived; it will use the same name and location as the Output feature class parameter but with a _Pnt suffix.
Output Feature Class Simplification Algorithm FOINT_REMOVE Simplification Tolerance		Specifies whether to create an output point feature class to store the centers of polygons that are removed because they are smaller than the Minimum area parameter. The point output is derived; it will use the same name and location as the Output feature class parameter but with a _Pnt suffix.
Output Feature Class Simplification Algorithm FOINT_REMOVE Simplification Tolerance Infinitum Area (optional) Handling Topological Errors (optional) RESOLVE_ENDORS Keep collapsed points (optional)		Specifies whether to create an output point feature class to store the centers of polygons that are removed because they are smaller than the Minimum area parameter. The point output is derived; it will use the same name and location as the Output feature class parameter but with a _Pnt suffix.
Output Feature Class Simplification Algorithm FOINT_REMOVE Simplification Tolerance		Specifies whether to create an output point feature class to store the centers of polygons that are removed because they are smaller than the Minimum area parameter. The point output is derived; it will use the same name and location as the Output feature class parameter but with a _Pnt suffix.
Uutput Feature Class Sisplification Algoriths FOINT_REMOVE Sisplification Tolerance		Specifies whether to create an output point feature class to store the centers of polygons that are removed because they are smaller than the Minimum area parameter. The point output is derived; it will use the same name and location as the Output feature class parameter but with a _Pnt suffix.

969 2. Calculating the total number of nodes using ArcGIS (Figure AFigure S22):

)	Add a new field in the attribute table of dataset.	
l	Open Field Calculator.	
2	Switch the parser to python mode, and enter the following code "!shape.pointcount!" in the b	<u>ue box</u>
3	to calculate the total number of nodes for each glacial lake boundary.	
	Field Calculator ×	
	Parser	
	○ VB Script	
	Fields: Type: Functions:	
	FID Conjugate()	
	Shape	
	OBJECTID .numerator()	
	IMGSOURCE O Date .real() NDWI_T .as_integer_ratio()	
	.fromhex()	
	Type	
	ACODATE math.acos()	
	Shape_Leng math.acosh() math.asin()	
	Show Codeblock	
	Check =	
	!shape.pointcount!	
	About calculating fields Clear Load Save	
	OK Cancel	
	Figure SA2 2. Total node calculation in ArcGIS.	
	<u>Agure BA22. Total houe calculation in Arcons.</u>	
I	3. Calculating the number of inner nodes:	
	For polygons without islands (Figure AFigure S33), use the equation 5. An inner node is a polygon	
	where the interior angle surrounding it is greater than 180 degrees. An outer node is the opposite	vertex
	inner node, where the interior angle is less than 180 degrees. We found that the outer nodes are usual	
	minor nous, where the interior angle is less than 100 degrees. We round that the outer nous are usual	of the
		of the
	more than the inner nodes in our glacial lake dataset. The total nodes in ArcGIS contain one over	of the lly four apping
	more than the inner nodes in our glacial lake dataset. The total nodes in ArcGIS contain one over node to close the polygon, meaning the endpoint is also the startpoint. This extra count was deleted	of the lly four apping
	more than the inner nodes in our glacial lake dataset. The total nodes in ArcGIS contain one over	of the lly four apping



🔨 Feature To Polygon			>	×
• Input Features			Feature To Polygon	,
polyline		- 🖻	Creates a feature class containing polygons generated from areas enclosed by input line or	
		+	polygon features.	
		×	LINE INPUT OUTPUT	
		Ť		
		4		
Output Feature Class new polygon		2	POLYGON INPUT OUTPUT	
XY Tolerance (optional)	Meters	~		
Preserve attributes (optional				
Label Features (optional)				
			LINE AND POLYGON INPUTS OUTPUT	
		0		
	OK Cancel Environments	<< Hide Help	Tool Help	
nany islands ther	e are in each lake (Figure		n, which outputs the islands. Then we can be seen as a second sec	<u> </u>
Erase			- L >	×
		^	– ⊔ >	×
Input Features		- 🖻	Erase Creates a feature class by overlaying the Input Features with the polygons of the Erase	×
Input Features new polygon Frase Features initial polygon		· 🖻 ^		×
Input Features new polygon Frase Features initial polygon		• 🖻	Erase Creates a feature class by overlaying the Input Features with the polygons of the Erase Features. Only those portions of the input features falling outside the erase features outside	×
Input Features new polygon Erass Features initial polygon Output Feature Class	Reters	• 🖻	Erase Creates a feature class by overlaying the Input Features with the polygons of the Erase Features. Only those portions of the input features falling outside the erase features outside boundaries are copied to the output feature class.	×
Input Features new polygon Erase Features initial polygon Output Feature Class polygon of island	Iters	• 🖻	Erase Creates a feature class by overlaying the Input Features with the polygons of the Erase Features. Only those portions of the input features falling outside the erase features outside	×
Input Features new polygon Frase Features initial polygon Output Feature Class polygon of island	Kters	• 🖻	Erase Creates a feature class by overlaying the Input Features with the polygons of the Erase Features. Only those portions of the input features falling outside the erase features outside boundaries are copied to the output feature class.	×
Input Features new polygon Frase Features initial polygon Output Feature Class polygon of island		• 🖻	Erase Creates a feature class by overlaying the Input Features with the polygons of the Erase Features. Only those portions of the input features falling outside the erase features outside boundaries are copied to the output feature class.	×
Input Features new polygon Frase Features initial polygon Output Feature Class polygon of island		• 🖻	Erase Creates a feature class by overlaying the Input Features with the polygons of the Erase Features. Only those portions of the input features falling outside the erase features outside boundaries are copied to the output feature class.	×
Input Features new polygon Frase Features initial polygon Output Feature Class polygon of island		• 🖻	Erase Creates a feature class by overlaying the Input Features with the polygons of the Erase Features. Only those portions of the input features falling outside the erase features outside boundaries are copied to the output feature class.	×
Input Features new polygon Frase Features initial polygon Output Feature Class polygon of island	Iters	• 🖻	Erase Creates a feature class by overlaying the Input Features with the polygons of the Erase Features. Only those portions of the input features falling outside the erase features outside boundaries are copied to the output feature class.	×
Input Features Rese Features Initial polygon Output Feature Class polygon of Island		• 🖻	Erase Creates a feature class by overlaying the Input Features with the polygons of the Erase Features. Only those portions of the input features falling outside the erase features outside boundaries are copied to the output feature class.	×
Input Features Rese Features Initial polygon Output Feature Class polygon of Island	Katars	• 🖻	Erase Creates a feature class by overlaying the Input Features with the polygons of the Erase Features. Only those portions of the input features falling outside the erase features outside boundaries are copied to the output feature class.	×
Input Features Rese Features Initial polygon Output Feature Class polygon of Island	Kters	• 🖻	Erase Creates a feature class by overlaying the Input Features with the polygons of the Erase Features. Only those portions of the input features falling outside the erase features outside boundaries are copied to the output feature class.	×
Input Features Rese Features Initial polygon Output Feature Class polygon of Island	Inters	• 🖻	Erase Creates a feature class by overlaying the Input Features with the polygons of the Erase Features. Only those portions of the input features falling outside the erase features outside boundaries are copied to the output feature class.	×
Input Features Rese Features Initial polygon Output Feature Class polygon of Island		• 🖻	Erase Creates a feature class by overlaying the Input Features with the polygons of the Erase Features. Only those portions of the input features falling outside the erase features outside boundaries are copied to the output feature class.	×
Input Features Rese Features Initial polygon Output Feature Class polygon of Island		• 🖻	Erase Creates a feature class by overlaying the Input Features with the polygons of the Erase Features. Only those portions of the input features falling outside the erase features outside boundaries are copied to the output feature class.	×
Input Features Rese Features Initial polygon Output Feature Class polygon of Island	Kters	• 🖻	Erase Creates a feature class by overlaying the Input Features with the polygons of the Erase Features. Only those portions of the input features falling outside the erase features outside boundaries are copied to the output feature class.	X
Input Features new polygon Frase Features initial polygon Output Feature Class polygon of island	Kters	• 🖻	Erase Creates a feature class by overlaying the Input Features with the polygons of the Erase Features. Only those portions of the input features falling outside the erase features outside boundaries are copied to the output feature class.	×
Input Features new polygon Frase Features initial polygon Output Feature Class polygon of island			Erase Creates a feature class by overlaying the Input Features with the polygons of the Erase Features. Only those portions of the input features failing outside the erase features outside boundaries are copied to the output feature class. INPUT ERASE FEATURE OUTPUT OUTPUT	×
Input Features new polygon trase Features initial polygon Output Feature Class polygon of Island XY Tolerance (optional)	OK Cancel Environments		Erase Creates a feature class by overlaying the Input Features with the polygons of the Erase Features. Only those portions of the input features failing outside the erase features outside boundaries are copied to the output feature class. INPUT ERASE FEATURE OUTPUT OUTPUT	×
b Input Festures new polygon b Inse Festures initial polygon Output Festure Class polygon of Island IY Tolerance (optional)			Erase Creates a feature class by overlaying the Input Features with the polygons of the Erase Features. Only those portions of the input features failing outside the erase features outside boundaries are copied to the output feature class. INPUT ERASE FEATURE OUTPUT OUTPUT	×
• Input Features new polygon • Irase Features initial polygon Output Feature Class polygon of Island XY Tolerance (optional) 	OK Cancel Environments		Erase Creates a feature class by overlaying the Input Features with the polygons of the Erase Features. Only those portions of the input features failing outside the erase features outside boundaries are copied to the output feature class. INPUT ERASE FEATURE OUTPUT OUTPUT	×
Figure SA77. Era	OK Cancel Environments se tool in ArcGIS.	Keiner (Keiner)	Erase Creates a feature class by overlaying the Input Features with the polygons of the Erase Features. Only those portions of the input feature falling outside the erase features outside boundaries are copied to the output feature class.	×
Erase Features initial polygon Output Feature Class polygon of island If Tolerance (optional) Figure SA77. Era	OK Cancel Environments se tool in ArcGIS.	Keiner (Keiner)	Erase Creates a feature class by overlaying the Input Features with the polygons of the Erase Features. Only those portions of the input features failing outside the erase features outside boundaries are copied to the output feature class. INPUT ERASE FEATURE OUTPUT OUTPUT	×
Figure SA77. Era	OK Environments se tool in ArcGIS. the number of inner node	 Key State Key State	Erase Creates a feature class by overlaying the Input Features with the polygons of the Erase Features. Only those portions of the input feature falling outside the erase features outside boundaries are copied to the output feature class. INPUT ERASE FEATURE OUTPUT OUTPUT Tool Help	×
Figure SA77. Era	OK Cancel Environments se tool in ArcGIS.	 Key State Key State	Erase Creates a feature class by overlaying the Input Features with the polygons of the Erase Features. Only those portions of the input feature falling outside the erase features outside boundaries are copied to the output feature class. INPUT ERASE FEATURE OUTPUT OUTPUT Tool Help	×