https://doi.org/10.5194/essd-2021-462
Preprint. Discussion started: 20 January 2022

Open Access Earth System Science Data Discussions

A New Operational Mediterranean Diurnal Optimally Interpolated SST Product within the Copernicus Marine Environment Monitoring Service

Andrea Pisano[1], Daniele Ciani[1], Salvatore Marullo[1,2] Rosalia Santoleri[1], Bruno Buongiorno Nardelli[3]

[1]CNR-ISMAR, Via del Fosso del Cavaliere 100, Rome, 00133, Rome, Italy
[2]ENEA, Via Enrico Fermi, 45, 00044 Frascati, Italy
[3]CNR-ISMAR, Calata Porta di Massa, Napoli, 80133 , Italy

Correspondence to: Andrea Pisano (andrea.pisano@cnr.it)

Abstract. Within the Copernicus Marine Environment Monitoring Service (CMEMS), a new operational MEDiterranean Diurnal Optimally Interpolated SST (MED DOISST) product has been developed. This product provides hourly mean maps (Level-4) of sub-skin SST at 1/16° horizontal resolution over the Mediterranean Sea from January 2019 to present. The product is built by combining hourly SST data from the Spinning Enhanced Visible and InfraRed Imager (SEVIRI) on board Meteosat Second Generation and model analyses through optimal interpolation. SEVIRI and model data are respectively used as the observation source and first-guess. The differences between satellite and model SST are free, or nearly free, of any diurnal cycle, thus allowing them to be interpolated in space and time using satellite data acquired at different times of the day. The accuracy of the MED DOISST product is assessed here by comparison against surface drifting buoy measurements, covering the years 2019 and 2020. The diurnal cycle reconstructed from DOISST is in good agreement with the one observed by independent drifter data, with a mean bias of 0.041 ± 0.001 K and root-mean-square difference (RMSD) of 0.412 ± 0.001 K. The new SST product is more accurate than the input model during the central warming hours, when the model, on average, underestimates drifter SST by one tenth of degree. The MED DOISST product is also able to reproduce accurately the extreme diurnal warming events frequently observed in the Mediterranean Sea, which may reach amplitudes larger than 5 K during the warm season. This product can contribute to improve the prediction capability of numerical weather forecast systems (e.g., through improved forcing/assimilation), as well as the monitoring of surface heat budget estimates and temperature extremes which can have significant impacts on the marine ecosystem.

The full MED DOISST product (released on 04 May 2021) is available upon free registration at https://doi.org/10.25423/CMCC/SST_MED_PHY_SUBSKIN_L4_NRT_010_036 (Pisano et al., 2021). The reduced subset used here for validation and review purposes is openly available at https://doi.org/10.5281/zenodo.5807729 (Pisano, 2021).

https://doi.org/10.5194/essd-2021-462
Preprint. Discussion started: 20 January 2022

1 Introduction

In the last decades, the development of accurate satellite-based Sea Surface Temperature (SST) products required an increasing effort to meet an ever-growing request from scientific, operational and emerging policy needs. Indeed, infrared and/or microwave satellite radiometers allow a systematic and synoptic mapping of the ocean surface temperature (under clear-sky conditions for the infrared and in the absence of rain for the microwave bands) with spatial resolutions from one to few kilometers and temporal sampling from hourly to daily (Minnett et al., 2019). This almost continuous coverage represents a unique characteristic of satellite thermal data, which is clearly not achievable with the use of in situ measurements alone. Indeed, though in situ sensors reach significantly higher accuracy than satellite sensors, with uncertainties that can reach O(10^{-2} °C), they provide pointwise seawater temperature measurements, generally characterized by a poor and non-uniform sampling of the ocean surface.

There is a huge variety of satellite-based SST datasets, characterized by different nominal resolutions as well as temporal and spatial (global or regional) coverage, and based on different processing algorithms and satellite sensors, but designed to provide highly accurate SST estimates (Yang et al., 2021). Operational datasets are typically distributed in near real time (NRT), delayed-mode or as reprocessed datasets, and may include different processing levels, from single satellite passes processed to provide valid SST values in the original observation geometry, the so-called Level-2 (L2), to images remapped onto a regular grid, also known as Level-3 (L3), up to the spatially complete Level-4 (L4), interpolated over fixed regular grids. These latter are required by several applications since the lower levels are typically affected by several data voids (due to clouds, rain, land, sea-ice, or other environmental factors depending on the type of sensors). The timely availability of SST data, ranging from a few hours to a few days before real time, allows their use as boundary condition and/or assimilation in meteorological and ocean forecasting systems (Waters et al., 2015), to improve the retrieval of ocean surface currents (Bowen et al., 2002; Rio and Santoleri 2018), and monitor some weather extreme events, such as marine heatwaves (Oliver et al., 2021). The reprocessing of long-term SST data records, typically covering the satellite era (1981-present), aims to provide more stable and consistent datasets, complementing the NRT production, to be used to investigate climate variability and monitor changes from interannual to multi-decadal timescales (Deser et al., 2010), including e.g. SST trends' estimates (Good et al., 2007; Pisano et al., 2020). The Copernicus Marine Environment Monitoring Service (CMEMS) is one of the main examples of how satellite observations, including not only SST but a wide range of surface variables (e.g., sea surface salinity, sea surface height, ocean color, winds and waves), are exploited to derive and disseminate high-level products (Le Traon et al., 2019), namely L4 data in order to be directly usable for downstream applications.

The majority of the existing L4 SST datasets are provided as daily, weekly or monthly averaged fields (see e.g. Fiedler et al., 2019; Yang et al., 2021). Examples of well-known state-of-the-art SST daily datasets include the Global Ocean Sea Surface

https://doi.org/10.5194/essd-2021-462
Preprint. Discussion started: 20 January 2022

Temperature and Sea Ice (OSTIA) dataset (Good et al., 2020), the European Space Agency (ESA) Climate Change Initiative (CCI), the Copernicus Climate Change Service (C3S) Reprocessed Sea Surface Temperature Analyses (Merchant et al., 2019), and the NOAA Daily Optimally Interpolated SST (OISST) v2.1 dataset, previously known/referred to as Reynolds SST analysis (Huang et al., 2021). Though a daily resolution is generally sufficient to meet the requirements of many of the oceanographic applications, it does not resolve the SST diurnal cycle, the typical day-night SST oscillation mainly driven by solar heating. Within the oceanic thermal skin layer (few µm to 1 mm), SST is typically subject to a large potential diurnal cycle (especially under low wind speed and strong solar heating conditions) reaching amplitudes up to 3 K in the world oceans (Gentemann et al., 2008; Gentemann and Minnett, 2008).

The SST diurnal cycle has several implications on mixed layer dynamics, air-sea interaction and the modulation of the lower atmosphere dynamics. The most direct consequence of the SST diurnal amplitude variability is certainly on air-sea fluxes. Clayson and Bogdanoff (2013) estimated that the diurnal SST cycle contributes with slightly less than that 5 Wm^{-2} to the global ocean-atmosphere heat budget with peaks of about 10 Wm^{-2} in the Tropics. The inclusion of a realistic diurnal SST cycle in atmospheric numerical simulation also has a non-negligible impact on cloud dynamics. Chen and Houze (1997) have shown that in the Tropical Warm Pool, where extreme localized warming events occur, the diurnal warming can contribute to modulate the evolution of convective clouds and, more in general, can impact the ocean-atmosphere coupling in numerical models, producing a more realistic spatial pattern of warming and precipitation (Bernie et al., 2008). Overall, the diurnal cycle of SST is generally underestimated in current ocean models and the assimilation of SST at high temporal frequency has the potential to improve sea surface variability and mixed layer accuracy (Storto and Oddo, 2019).

In principle, the best opportunity to measure the diurnal cycle comes from infrared radiometers on board geostationary satellites. Their observations are sufficiently accurate and frequent to resolve the diurnal signal variability whenever cloud cover is not too persistent. An example is provided by the Spinning Enhanced Visible Infra-Red Imager (SEVIRI) onboard the Meteosat Second Generation (MSG) geostationary satellite covers. The operational retrieval of SST from MSG/SEVIRI (managed by the European Organization for the Exploitation of Meteorological Satellites, EUMETSAT, Ocean and Sea-Ice Facility, OSI-SAF) produces L3C hourly sub-skin SST products by aggregating 15 minutes (MSG/SEVIRI) observations within 1 hour. The sub-skin SST is the temperature at the base of the conductive laminar sub-layer of the ocean surface, as defined by the Group of High Resolution SST (GHRSST, see e.g. Minnett et al., 2019). In practice, this is the temperature at ~1 mm depth, and thus particularly sensitive to diurnal warming.

For the global ocean, the Operational Sea surface Temperature and sea Ice Analysis (OSTIA) diurnal product (While et al., 2017) provides daily gap-free maps of hourly mean skin SST at 0.25° x 0.25° horizontal nominal resolution, using in situ and satellite data from infrared radiometers. The skin SST is the temperature within the conductive diffusion-dominated sub-layer at a depth of ~10-20 µm (as defined by GHRSST, Minnett et al., 2019). This system produces a skin SST by combining the

https://doi.org/10.5194/essd-2021-462
Preprint. Discussion started: 20 January 2022

OSTIA foundation SST analysis (Good et al., 2020) with a diurnal warm-layer temperature difference and a cool skin temperature difference derived from numerical models.

At regional scale, a method to reconstruct the hourly SST field over the Mediterranean Sea from SEVIRI data has been proposed by Marullo et al. (2014, 2016). The reconstruction is based on a blending of satellite observations and numerical model analyses (used as first-guess) using optimal interpolation. Though model analyses by definition also assimilate observations, which could thus in principle include hourly SEVIRI data, in the present configuration they are not able to deal with such frequent updates (see section 2.2), and the approach presented here represents an effective way to improve the reconstruction of SST daily cycle from high-repetition satellite measurements. Previous works demonstrated the capability of SEVIRI to resolve the SST diurnal variability and to reconstruct accurate L4 SST hourly fields over the Mediterranean Sea, a basin that exhibits large diurnal SST variations (Buongiorno Nardelli et al., 2005; Minnett et al., 2019) that can easily exceed extreme values (~5 K) as observed in the Tropical Pacific (Chen and Houze 1997), in the Atlantic Ocean and other marginal seas (Gentemann et al., 2008; Merchant et al., 2008). The aim of this paper is to describe the operational implementation of a diurnal optimally interpolated SST (DOISST) product for the Mediterranean Sea (MED) at 1/16° horizontal resolution, building on the algorithm by Marullo et al. (2014, 2016). The assessment of the MED DOISST product covers two complete years (2019-2020), thus extending previous similar validations (Marullo et al., 2016).

2 The data

2.1 Satellite data

Input satellite SST is derived from the SEVIRI sensor onboard the Meteosat Second Generation (Meteosat-11) satellite. SEVIRI has a repeat cycle of 15 minutes over the 60S-60N and 60W-60E domain: Atlantic Ocean, European Seas and western Indian Ocean. The retrieval of SST from Meteosat-11/SEVIRI is managed by EUMETSAT OSI-SAF, which provides sub-skin SST data as aggregated (L3C) hourly products remapped onto a 0.05° regular grid. Hourly products result from compositing the best SST measurements available in one hour and are made available in near real time with a timeliness of 3 hours (see the OSI-SAF product user manual, https://osi-saf.eumetsat.int/products/osi-206). File format follows the Data Specification (GDS) version 2 from the Group for High Resolution Sea Surface Temperatures (GHRSST, https://podaac-tools.jpl.nasa.gov/drive/files/OceanTemperature/ghrsst/docs/GDS20r5.pdf). The computation of SST in day and night conditions is based on a nonlinear split window algorithm whose coefficients are determined from brightness temperature simulations on a radiosonde profile database, with an offset coefficient corrected relative to buoy measurements. A correction term derived from simulated brightness temperatures with an atmospheric radiative transfer model is then applied to the multispectral derived SST (OSI-SAF PUM, https://osi-saf.eumetsat.int/lml/doc/osisaf_cdop3_ss1_pum_geo_sst.pdf). L3C data are provided with additional information, including quality level and cloud flags. Such quality flags are provided at pixel level, ranging over a scale of five levels with increasing reliability: 1 (="cloudy"), 2 (="bad"), 3 (="acceptable"), 4 (="good") to 5 (="excellent").

The accuracy of Meteosat-11 SST data has been assessed through comparison with co-located drifting buoys, for day and night data separately covering the period from February to June 2018 (see the OSI-SAF scientific validation report, https://osi-saf.eumetsat.int/lml/doc/osisaf_cdop2_ss1_geo_sst_val_rep.pdf). The mean bias and standard deviation (derived from the differences between SEVIRI SSTs and drifter measurements over a matchup database) during nighttime have been quantified in -0.1 K and 0.53 K, respectively. During daytime, the bias remains practically unchanged (-0.09 K) and the standard deviation slightly higher (0.56 K). These statistics were derived by selecting SEVIRI SST with quality flags ≥ 3, and it is shown that the quality of SST improves when choosing higher quality levels.

For our purposes, we selected L3C SST data with quality flag ≥ 3, as also indicated/suggested in the OSI-SAF scientific validation report.

2.2 Model data

The model output fields of surface temperature are derived from the CMEMS Mediterranean Sea Physical Analysis and Forecasting product, and identified as MEDSEA_ANALYSIS_FORECAST_PHY_006_013 (https://resources.marine.copernicus.eu/product-detail/MEDSEA_ANALYSISFORECAST_PHY_006_013/INFORMATION; https://doi.org/10.25423/CMCC/MEDSEA_ANALYSISFORECAST_PHY_006_013_EAS6; last access: 03 November 2021; Clementi et al., 2021), and routinely produced by the CMEMS Mediterranean Monitoring and Forecasting Center (Med-MFC). The modelling system is based on the Mediterranean Forecasting System, MFS (Pinardi et al., 2003), a coupled hydrodynamic-wave model implemented over the Mediterranean basin, extended into the Atlantic Sea in order to better resolve the exchanges with the Atlantic Ocean at the Strait of Gibraltar, with a horizontal grid resolution of 1/24˚ (~4 km) and 141 unevenly spaced vertical levels (Clementi et al., 2017). The Ocean General Circulation Model is based on the Nucleus for European Modelling of the Ocean (NEMO v3.6) (Oddo et al., 2014, 2009), while the wave component is provided by Wave Watch-III. The model solutions are corrected by a variational data assimilation scheme (3DVAR) of temperature and salinity vertical profiles and along track satellite sea level anomaly observations (Dobricic and Pinardi 2008). The CMEMS Mediterranean SST L4 product (CMEMS product reference: SST_MED_SST_L4_NRT_OBSERVATIONS_010_004, https://resources.marine.copernicus.eu/product-detail/SST_MED_SST_L4_NRT_OBSERVATIONS_010_004/INFORMATION; last access: 03 November 2021) is used for the correction of surface heat fluxes with the relaxation constant of 110 $Wm^{-2}K^{-1}$ centered at midnight since the product provides foundation SST (~SST at midnight).

https://doi.org/10.5194/essd-2021-462
Preprint. Discussion started: 20 January 2022

The Med-MFC product is produced with two different cycles: a daily cycle for the production of forecasts (i.e., ten-days forecast on a daily basis), and a weekly cycle for the production of analyses. For our purposes, only hourly mean fields of sea surface temperature, which correspond to the first vertical level of the model centered at ~1 m from the surface, are selected.

2.3 In situ data

In situ data have been used for validation purposes (Section 4). Specifically, only surface drifting buoys have been used due to both their closeness to the sea surface (typically ~20 cm from the surface; Reverdin et al., 2010) and to their much larger number compared to other in situ instruments, which allows a more consistent and homogeneous temporal and spatial coverage.

Drifter data have been obtained from the CMEMS IN SITU (INS) TAC (identified as INSITU_MED_NRT_OBSERVATIONS_013_035, https://resources.marine.copernicus.eu/product-detail/INSITU_MED_NRT_OBSERVATIONS_013_035/INFORMATION; and INSITU_IBI_NRT_OBSERVATIONS_013_033, https://resources.marine.copernicus.eu/product-detail/INSITU_IBI_NRT_OBSERVATIONS_013_033/INFORMATION; last access: 03 November 2021), which collects and distributes a variety of physical and biogeochemical seawater measurements, provided with the same homogeneous file format . Each in situ measurement, including drifters, undergoes automated quality controls before its distribution. The quality of the data is expressed by control flags indexed from 0 to 9, with the value of 1 indicating best quality. Drifter data have been used to compile an hourly matchup database of co-located (in space and time) diurnal optimally interpolated SST (DOISST) values and model outputs (Section 4.1), and validation statistics are based on the comparison between DOISST, model SST and drifting buoy measurements over the matchup database (Section 4.2).

3 The Mediterranean diurnal optimally interpolated SST product

3.1 Product overview

The Mediterranean diurnal optimally interpolated SST (hereafter referred to as MED DOISST) operational product consists of hourly mean gap-free (L4) satellite-based estimates of the sub-skin SST over the Mediterranean Sea (plus the adjacent Eastern Atlantic box, see Section 2.2) at 0.0625° x 0.0625° grid resolution, from 1st January 2019 to near real time. Specifically, the product is updated daily and provides 24 hourly mean data of the previous day, centered at 00:00, 01:00, 02:00,…, 23:00 UTC. The MED DOISST product is published on the CMEMS on line catalogue and identified as SST_MED_PHY_SUBSKIN_L4_NRT_010_036 (CMEMS product reference) and cmems_obs-sst_med_phy-sst_nrt_diurnal-oi-0.0625deg_PT1H-m (CMEMS dataset reference). A synthesis of the product characteristics is shown Table 1.

CMEMS Product ID: SST_MED_PHY_SUBSKIN_L4_NRT_010_036

https://doi.org/10.5194/essd-2021-462
Preprint. Discussion started: 20 January 2022

Earth System Science Data
Open Access Discussions

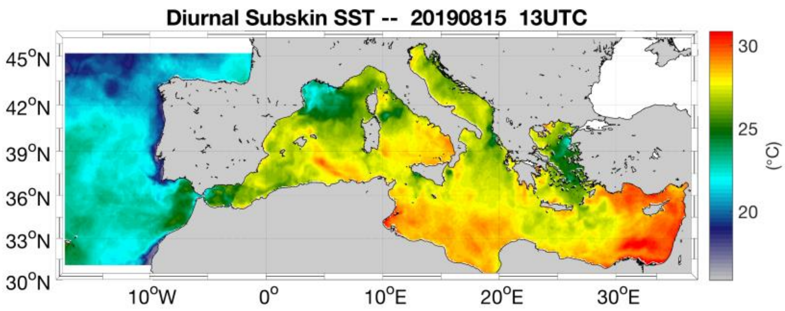
CMEMS Dataset ID: cmems_obs-sst_med_phy-sst_nrt_diurnal-oi-0.0625deg_PT1H-m	
General description	The CMEMS Mediterranean diurnal product provides near-real-time, hourly mean, gap-free (L4) sub-skin SST fields over the Mediterranean Sea and the adjacent Atlantic box over a 0.0625°x0.0625° regular grid, covering the period from 2019 to present (one day before real time). This product is built from optimal interpolating the Level-3C (merged single-sensor, L3C) SEVIRI data as observations and the CMEMS Mediterranean model analyses as first-guess. 
Horizontal resolution	0.0625° x 0.0625° (1/16°) degrees [871x253]
Temporal resolution	Hourly
Spatial coverage	Mediterranean Sea + adjacent North Atlantic box (W=-18.1250, E=36.2500, S=30.2500, N=46.0000)
Temporal coverage	2019/01/01 – near real time (-14H)
Variables	Sub-skin SST (K) Analysis Error (%)
Format	NetCDF – CF-1.4 convention compliant
DOI	https://doi.org/10.25423/CMCC/SST_MED_PHY_SUBSKIN_L4_NRT_010_036
Comments	Eventual updates of this product will be described in the corresponding Product User Manual (PUM) and Quality Information Document (QUID) available on the CMEMS on line catalogue.

Table 1. The CMEMS MED DOISST product description synthesis.

3.2 Processing chain

The system implements the DOISST scheme developed by Marullo et al. (2014). The system ingests merged single-sensor (L3C) SEVIRI data as the observation source, and the CMEMS Mediterranean Sea model outputs as first-guess. It has been shown that the diurnal signal in the hourly anomaly SST field (satellite-model) is reduced by about one order of magnitude with respect to the full signal, thus allowing to interpolate SST anomalies using satellite data acquired at different times of the

https://doi.org/10.5194/essd-2021-462
Preprint. Discussion started: 20 January 2022

day (Marullo et al., 2014). Several trials over a large variety of environmental conditions have shown that the temporal window to be used for the selection of input observations is ±24 hours.

The data sub-sampling strategy, inversion technique and numerical implementation of the optimal interpolation scheme are based on the CMEMS NRT MED SST processing chain (Buongiorno Nardelli et al., 2013), which provides daily mean fields of foundation SST over the Mediterranean Sea (CMEMS product reference: SST_MED_SST_L4_NRT_OBSERVATIONS_010_004, https://resources.marine.copernicus.eu/product-detail/SST_MED_SST_L4_NRT_OBSERVATIONS_010_004/INFORMATION; last access: 03 November 2021). Here, the diurnal SST chain is organized in three main modules (Fig. 1).

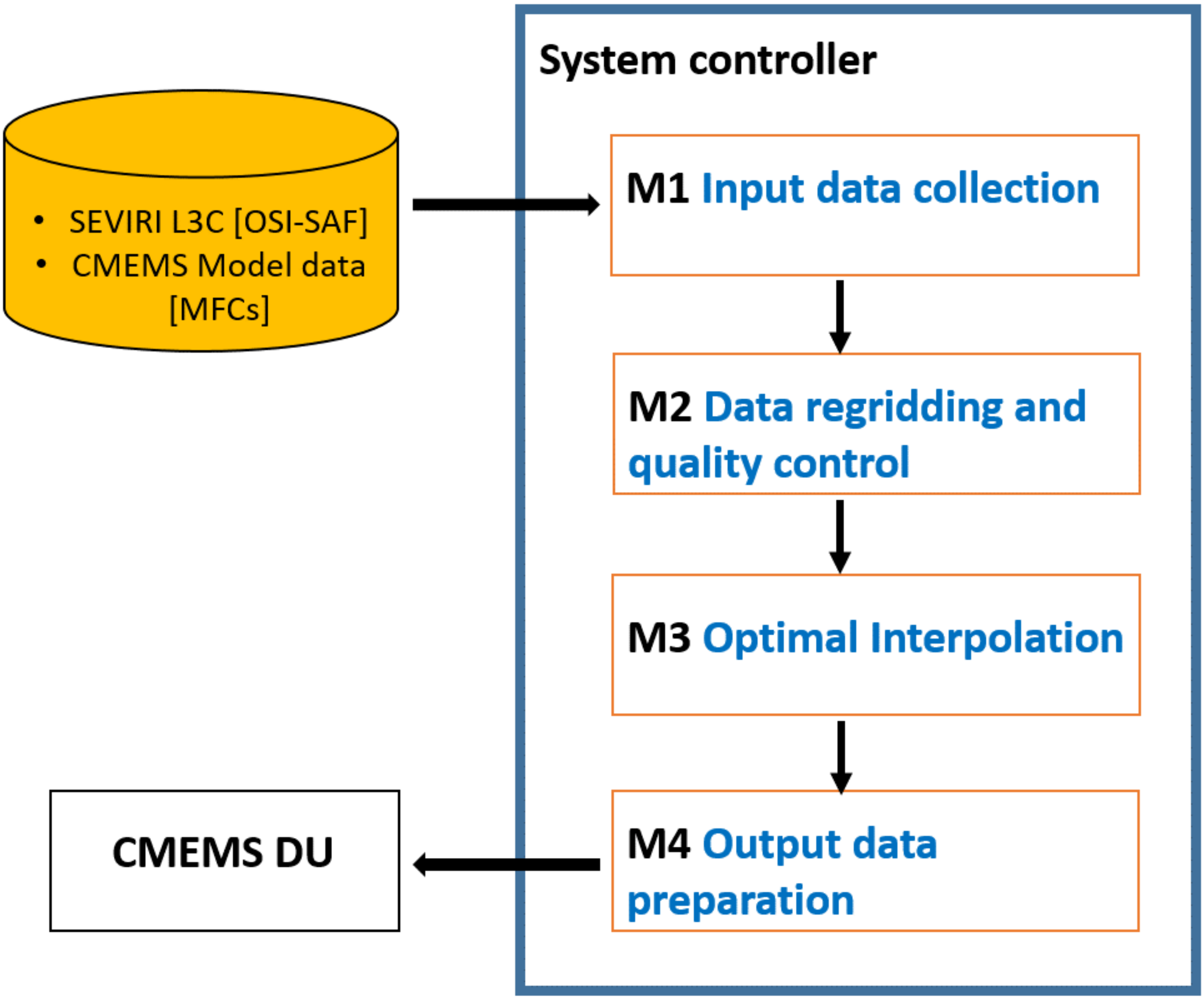


Figure 1. Schematic diagram of the processing chain used for the MED DOISST SST product.

https://doi.org/10.5194/essd-2021-462
Preprint. Discussion started: 20 January 2022

Module M1 manages the external interfaces to get both upstream L3C SST and model data: hourly mean L3C sub-skin SST data at 0.05° grid resolution are downloaded from OSI-SAF; hourly seawater potential temperatures at 1.0182 meter are obtained from the CMEMS Mediterranean Sea model outputs.

Module M2 extracts and regrids (through bilinear interpolation) L3C data and model outputs over the CMEMS Mediterranean Sea geographical area. A selection over SEVIRI is performed by flagging the pixels with quality flag < 3.

Module M3 performs a space-time optimal interpolation (OI) algorithm. L4 data are obtained as a linear combination of the SST anomalies, weighted directly with their correlation to the interpolation point and inversely with their cross-correlation and error. Correlations are typically expressed through analytical functions with predefined spatial and temporal de-correlation lengths. Here, the covariance function $f(r, \Delta t)$ is the one defined in Marullo et al. (2014), and given as the product of a spatial and temporal component:

$$f(r, \Delta t) = f(r) * f(\Delta t) = [a*\exp(-r/R) + (1.-a)/(1.+r)c] * \exp(- (\Delta t/T)d)$$

where r is the distance (in km) between the observation and the interpolation point; Δt is the temporal difference (in hours) between the observation and the interpolation point; R = 200 km is the decorrelation spatial length; T = 36 h is the decorrelation time length; the other parameters are set as follows: a = 0.70, c = 0.26, d = 0.4. All these parameters have been deduced from a statistical analysis of the satellite SST data.

The input data are selected only within a limited sub-domain (within a given space-time interval), with a temporal window of ±24 h (Marullo et al., 2014) and a spatial search radius of about 700 km (Buongiorno Nardelli et al., 2013). A check to avoid data propagation across land is performed between each pixel within the sub-domain and the given interpolation point (eventually discarded if there are land pixels between the straight line connecting the two points).

The interpolation error (analysis_error field in the L4 file, Table 1) is obtained from the formal definition of the error variance derived from optimal interpolation theory (e.g., Bretherton et al., 1976). This error ranges between [0,100%], meaning that all observations are used (no first-guess data are used) when the error is zero, while only first-guess data are used (i.e no observations available) when the error is 100%.

The optimal interpolation algorithm is synthetized as follows:

- Hourly SEVIRI and model SSTs in a space/time window of 700 km/ ±24 h around the interpolation position/time are ingested;
- SEVIRI data with quality flag ≥ 3 are retained;

- Regridding over the Mediterranean Sea;
- Hourly model SSTs are subtracted from valid SSTs producing SST anomalies;
- SST anomalies are used as data input for the optimal interpolation analysis;
- Optimal interpolation is run using the covariance function defined above;
- The model SST is added to the optimally interpolated output again.

The only difference with the original method is that all the input observations are interpolated, while in Marullo et al. (2014) valid SST observations are left unchanged (not interpolated).

4 Validation of diurnal product

4.1 Validation framework

The accuracy of the MED DOISST product has been assessed through comparison with independent co-located (in space and time) surface drifting buoy data (matchups). The validation framework is based on the compilation of a matchup database between DOISST and drifters measurements covering the full years 2019 and 2020. The large number of drifters provides a rather homogeneous and continuous spatial and temporal coverage over the whole period (Fig. 2) allowing a robust statistical approach.

Firstly, a pre-selection of high-quality drifter data is performed, retaining only temperatures with quality flag equal to 1 (good) or 2 (probably good) (see section 2.3). Then, the validation is carried out on hourly basis, building a matchup database by collecting the closest (in space) SST grid point to the in situ measurement within a symmetric temporal window of 30 minutes with respect to the beginning of each hour. A final quality control iteratively identifies drifter temperatures for which the module of the difference between satellite and drifter temperature exceeds n-times the standard deviation σ of the distribution of all these differences (δ). At each step of decreasing n, data that falls out of the interval $I = [mean(\delta) - n \cdot \sigma, mean(\delta) + n \cdot \sigma]$ are flagged as outliers and then not included in the next step. The process starts for n=10 and stops at n=3. This last quality control removes ~1% of the total original sampling (as expected from a gaussian distribution) of drifter data that clearly revealed anomalous temperature values.

Validation statistics are quantified in terms of mean bias and Root-Mean-Square Difference (RMSD) from matchup temperature differences (namely, SST minus drifter). Each statistical parameter is associated with a 95% confidence interval computed through a bootstrap procedure (Efron 1994).
In order to evaluate the DOISST performance with respect to the model, the same validation procedure has been applied to the modeled SST.

https://doi.org/10.5194/essd-2021-462
Preprint. Discussion started: 20 January 2022

4.2 Comparison with drifters

4.2.1 The mean diurnal cycle

The spatial distribution of DOISST and drifter matchups over the 2019-2020 period, along with their pointwise bias (i.e., DOISST minus drifter measurement) shows a rather homogeneous coverage over the most of the CMEMS MED domain (Fig. 2), although some areas are characterized by quite low coverage, such as the North Adriatic Sea or North Aegean Sea. The spatial distribution also evidences a positive tendency of the bias indicating that, on average, DOISSTs are warmer than drifters' temperatures.

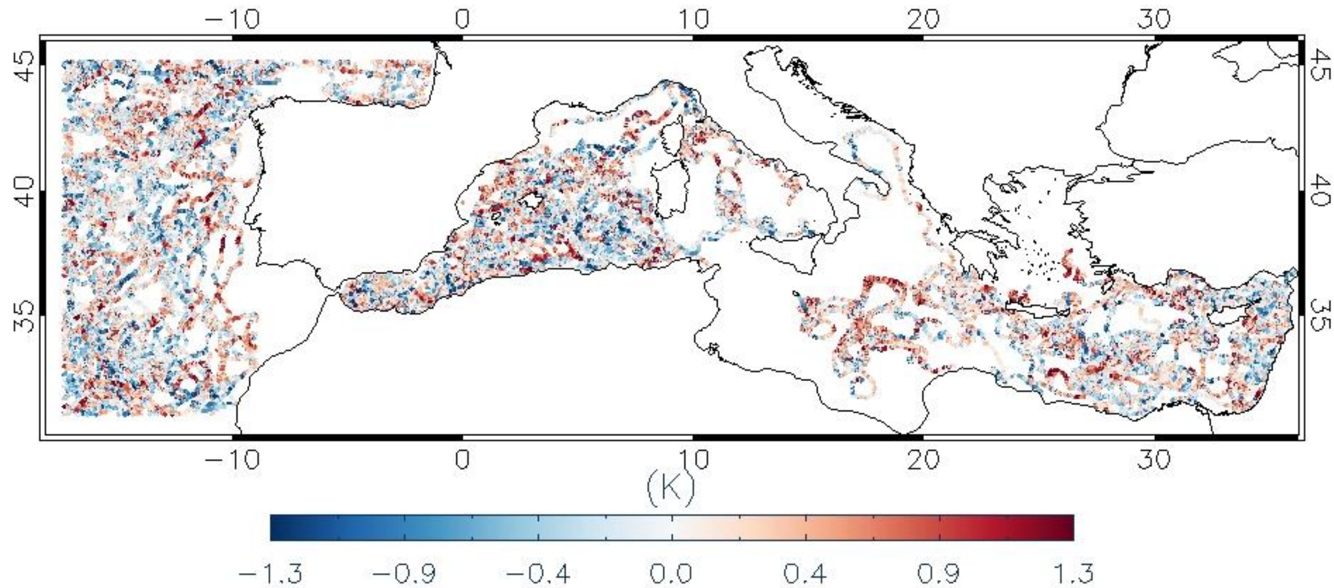


Figure 2. Spatial distribution of the matchup points along with their punctual bias (i.e., SST minus drifter data, K) over the CMEMS Mediterranean domain from 2019/01/01 to 2020/12/31.

The DOISST product shows effectively an overall small positive mean bias of 0.041 ± 0.001 K and a RMSD of 0.412 ± 0.001 K (Table 2). A negative bias of -0.100 ± 0.001 K and slightly larger RMSD of 0.467 ± 0.001 K characterize model SSTs.

	Period	Mean bias (K)	RMSD (K)	Matchups
DOISST	2019-01-01 to 2020-12-31	0.041 ± 0.001	0.412 ± 0.001	548959
Model	2019-01-01 to 2020-12-31	-0.100 ± 0.001	0.467 ± 0.001	548959

https://doi.org/10.5194/essd-2021-462
Preprint. Discussion started: 20 January 2022

Table 2. Summary statistics of DOISST and model outputs. Mean bias (K) and RMSD (K) are derived from temperature differences against drifters' data over the period 2019-2020.

The hourly mean bias of DOISST and model shows similar but opposite behaviour (Fig. 3, and Table 3). In both cases, the bias clearly exhibits a diurnal oscillation during the 24 hours but, while the bias of DOISST increases positively during the central diurnal warming hours, the one of the model increases negatively. The DOISST mean bias is practically null between 17:00 to 06:00 local time, ranging between -0.001 and 0.03 K, and highest (~0.1 K) between 10:00 and 13:00 local time. The bias of the model oscillates around ~-0.07 K between 23:00 and 07:00 local time. Then, it increases (in absolute value) reaching the peak of ~-0.16 K between 11:00 and 14:00 and decreases successively. Similar results are obtained for the RMSD, which increases with diurnal warming (Fig. 3, Table 3). However, the RMSD of DOISST is less impacted by diurnal variations, characterized by an amplitude of ~0.04 K against ~0.14 K of the model.

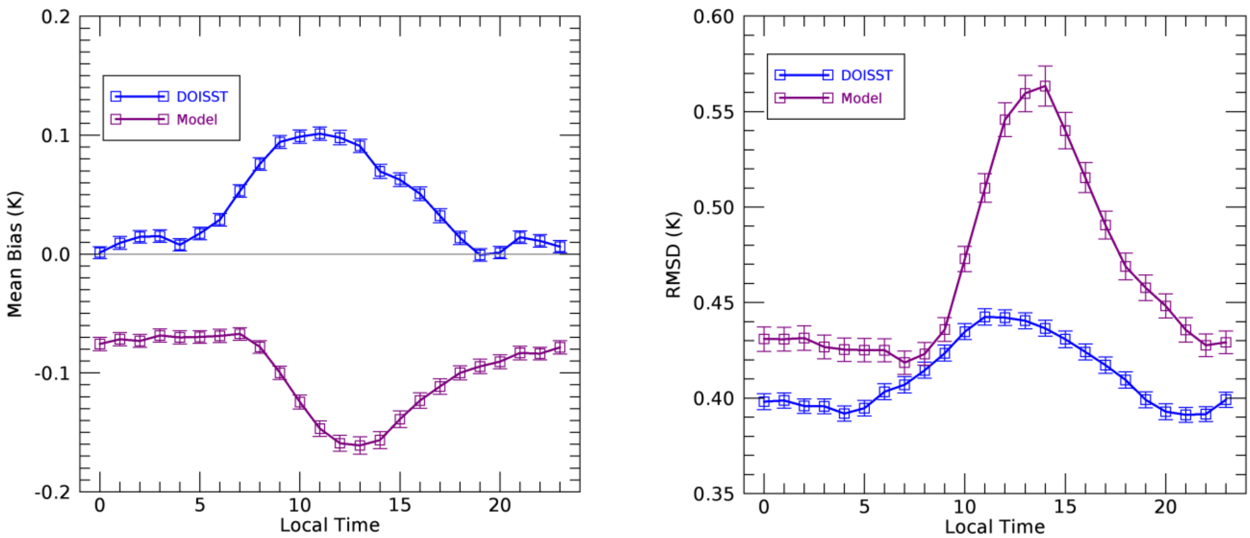


Figure 3. Mean bias (K) and RMSD (K) relative to MED DOISST (blue line) and model (purple line) based on the differences against drifters' data. Mean bias and RMSD are given as hourly mean over the period 2019-2020.

https://doi.org/10.5194/essd-2021-462
Preprint. Discussion started: 20 January 2022

Hour (local time)	Mean BIAS (K) (DOISST)	RMSD (K) (DOISST)	BUOY-AVAIL	Mean BIAS (K) (Model)	RMSD (K) (Model)
HH: 00	0.001 ± 0.005	0.398 ± 0.004	22807	-0.076 ± 0.006	0.431 ± 0.006
HH: 01	0.009 ± 0.005	0.399 ± 0.004	23004	-0.072 ± 0.006	0.431 ± 0.006
HH: 02	0.014 ± 0.005	0.396 ± 0.004	22798	-0.073 ± 0.005	0.431 ± 0.006
HH: 03	0.015 ± 0.005	0.396 ± 0.004	23078	-0.068 ± 0.006	0.427 ± 0.006
HH: 04	0.008 ± 0.005	0.392 ± 0.004	22857	-0.070 ± 0.005	0.425 ± 0.006
HH: 05	0.017 ± 0.005	0.395 ± 0.004	22806	-0.070 ± 0.005	0.425 ± 0.006
HH: 06	0.029 ± 0.005	0.403 ± 0.004	22819	-0.069 ± 0.006	0.425 ± 0.006
HH: 07	0.053 ± 0.005	0.407 ± 0.004	23379	-0.067 ± 0.005	0.419 ± 0.006
HH: 08	0.076 ± 0.005	0.415 ± 0.004	23501	-0.078 ± 0.006	0.423 ± 0.006
HH: 09	0.094 ± 0.005	0.423 ± 0.004	23481	-0.100 ± 0.006	0.436 ± 0.006
HH: 10	0.099 ± 0.006	0.435 ± 0.004	23270	-0.125 ± 0.006	0.473 ± 0.007
HH: 11	0.101 ± 0.006	0.442 ± 0.004	23311	-0.147 ± 0.006	0.510 ± 0.007
HH: 12	0.098 ± 0.006	0.442 ± 0.004	23129	-0.159 ± 0.007	0.546 ± 0.009
HH: 13	0.091 ± 0.006	0.440 ± 0.005	22836	-0.161 ± 0.007	0.560 ± 0.009
HH: 14	0.070 ± 0.006	0.436 ± 0.004	22673	-0.157 ± 0.007	0.563 ± 0.011
HH: 15	0.062 ± 0.006	0.431 ± 0.004	22418	-0.139 ± 0.007	0.540 ± 0.009
HH: 16	0.051 ± 0.006	0.424 ± 0.004	22368	-0.123 ± 0.007	0.515 ± 0.008
HH: 17	0.032 ± 0.006	0.417 ± 0.004	22019	-0.111 ± 0.006	0.491 ± 0.007
HH: 18	0.014 ± 0.006	0.410 ± 0.004	21916	-0.100 ± 0.006	0.469 ± 0.007
HH: 19	-0.001 ± 0.005	0.399 ± 0.004	22117	-0.095 ± 0.006	0.458 ± 0.007
HH: 20	0.001 ± 0.005	0.393 ± 0.004	22458	-0.090 ± 0.006	0.448 ± 0.006
HH: 21	0.014 ± 0.005	0.391 ± 0.004	23229	-0.083 ± 0.005	0.436 ± 0.006
HH: 22	0.011 ± 0.005	0.392 ± 0.004	23272	-0.084 ± 0.006	0.428 ± 0.006
HH: 23	0.006 ± 0.005	0.399 ± 0.004	23413	-0.078 ± 0.006	0.429 ± 0.006

Table 3. Summary statistics of MED DOISST and model products based on the differences against drifters' data over the matchup points. Mean bias (K), RMSD (K) and number of matchups are given as hourly mean over the period 2019-2020.

The mean diurnal cycle of DOISST (namely, the 24-hour mean SSTs estimated over the matchup dataset) is in very good agreement, within the error confidence interval, with the SST cycle reconstructed from drifters (Fig. 4). The two diurnal cycles are practically coincident between 17:00 and 06:00, while they are biased by ~0.1 K between sunrise and 16:00, coherently with the DOISST bias oscillation (Fig. 3). This bias could be related to skin SST getting warmer faster than 20 cm temperature. The diurnal cycle of model SST maintains always below that of in situ temperatures, evidencing larger differences during the central diurnal warming hours (Fig. 4). However, apart from the biases likely induced by the different depths, the SST

https://doi.org/10.5194/essd-2021-462
Preprint. Discussion started: 20 January 2022

amplitude as estimated from the DOISST and the model is ~2.3% larger and ~16% smaller than that of drifters, respectively, suggesting that the model tends to underestimate diurnal variations.

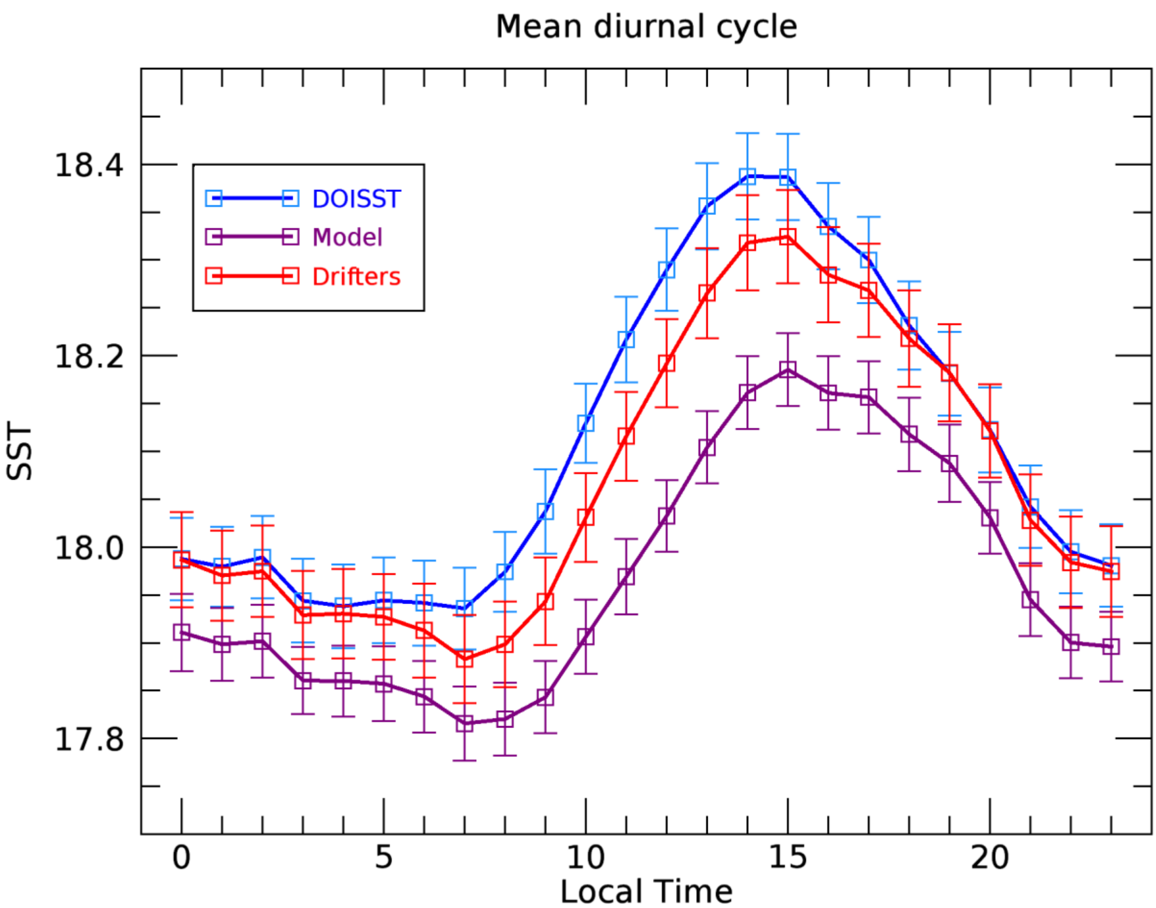


Figure 4. Mean diurnal cycle for MED DOISST (blue line), model (purple line) and drifters (red line) computed over the matchups from 2019 to 2020.

A delay of ~1 hour of the model with respect to DOISST and in situ on the onset of diurnal warming and in reaching the maximum is also evident. This delay could be explained as the physical result of delayed solar heating of the skin layer sensed by the satellite and of the first model layer. This may also be a consequence of the different packaging of the SEVIRI and model SST data into the hourly files: model hourly SST fields are centered at half of every hour (e.g., 12:30), while SEVIRI L3C at the beginning of each hour (e.g., 12:00) and obtained from collating data within one hour (from 11.30 to 12:29).

The capability of DOISST to capture and realistically reproduce diurnal variability is further investigated by analysing the seasonally averaged SST diurnal cycle (Fig. 5), computed as for the mean diurnal cycle (by using the matchup dataset) but over seasons: winter (December to February, D-J-F), spring (March to May, M-A-M), summer (June to August, J-J-A) and autumn (September to November, S-O-N). The effect of warming in the diurnal SST excursion is clearly more pronounced during spring and summer than winter and autumn, and reconstructed well in DOISST. During the warmer seasons, the DOISST shows the lower biases (Table 4), estimated in 0.036 ± 0.001 K (spring) and 0.012 ± 0.003 (summer). Conversely, the model reaches its higher biases, namely -0.101 ± 0.001 K (spring) and -0.117 ± 0.003 K (summer). The good agreement between DOISST and drifters during winter and autumn (Table 4) reveals that the hourly DOISST fields are reconstructed accurately also under cloudy conditions, which are more frequent during these seasons (Kotsias and Lolis, 2018).

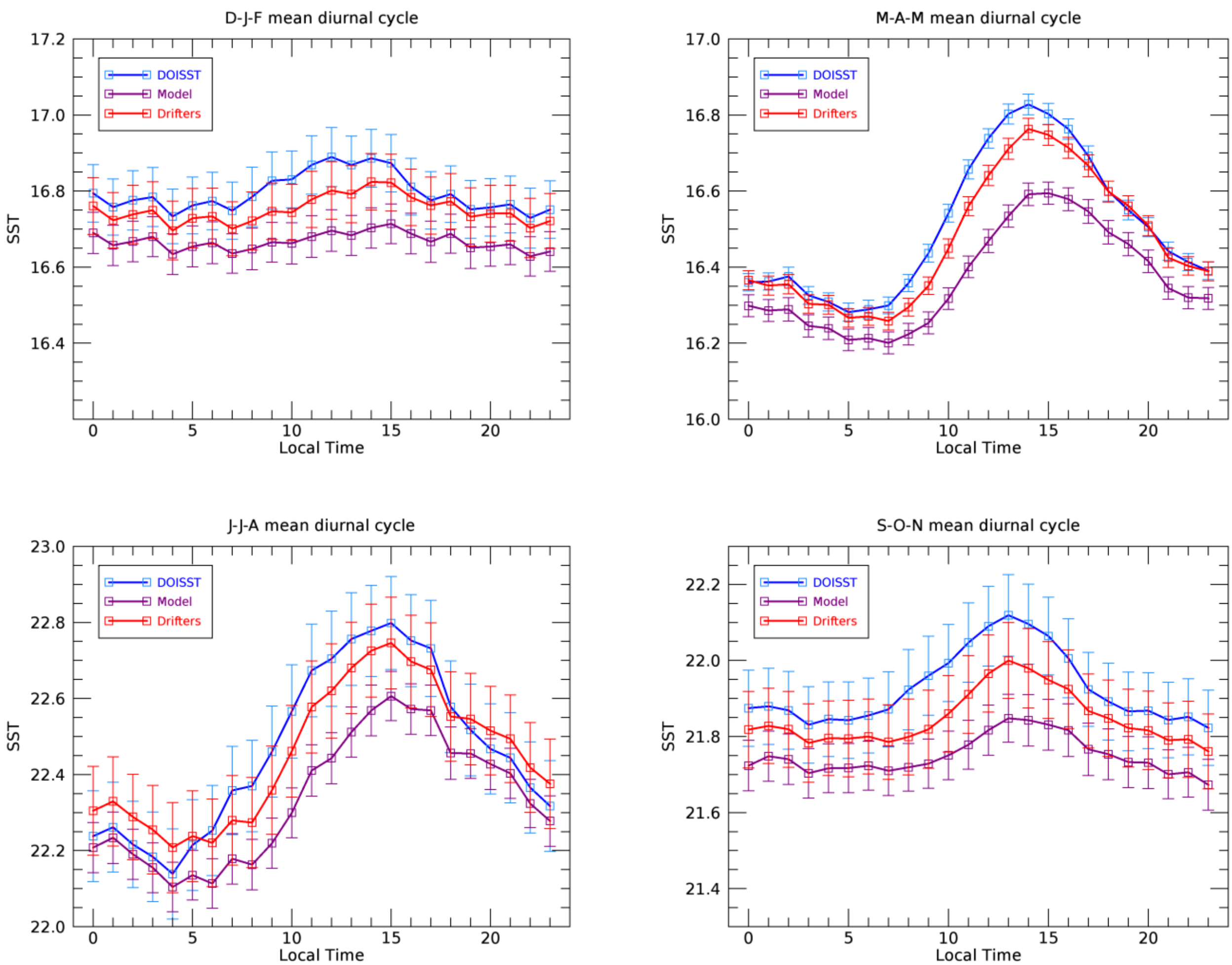


Figure 5. Seasonal mean diurnal cycle over the period 2019-2020 for MED DOISST (blue line), model (purple line) in situ (red line).

https://doi.org/10.5194/essd-2021-462
Preprint. Discussion started: 20 January 2022

	Period	Mean bias (K)	RMSD (K)	Matchups
D-J-F	DOISST	0.045 ± 0.003	0.428 ± 0.002	90247
	Model	-0.084 ± 0.004	0.563 ± 0.003	
M-A-M	DOISST	0.036 ± 0.001	0.383 ± 0.001	308448
	Model	-0.101 ± 0.001	0.389 ± 0.002	
J-J-A	DOISST	0.012 ± 0.003	0.483 ± 0.002	74107
	Model	-0.117 ± 0.003	0.486 ± 0.004	
S-O-N	DOISST	0.079 ± 0.003	0.429 ± 0.002	76157
	Model	-0.098 ± 0.004	0.590 ± 0.004	

Table 4. Summary statistics of DOISST and model outputs. Mean bias (K) and RMSD (K) are derived from temperature differences against drifters' data during winter (D-J-F), spring (M-A-M), summer (J-J-A) and autumn (S-O-N) over the period 2019-2020.

The capability of DOISST to reproduce diurnal warming events is analysed in the following section.

4.2.2 Diurnal warming events

Diurnal warming (DW) can be defined as the (positive) difference between the SST at a given time of the day and the foundation SST (see e.g. Minnett et al., 2019), i.e. the water temperature at a depth such that the daily variability induced by the solar irradiance is negligible. In many cases, the foundation SST coincides with the night minimum SST, namely the temperature that is recorded just before sunrise.

The capability of DOISST to describe diurnal warming events is analysed here in comparison with SEVIRI L3C, model and drifter data. The evaluation is carried out by computing daily DWAs from drifters and building a matchup dataset of DWAs as estimated from DOISST, SEVIRI L3C and model data. The inclusion of SEVIRI data is aimed at evaluating the impact of optimal interpolation on the input SEVIRI SSTs. The diurnal warming amplitude (DWA) is estimated here as a difference between the maximum occurred during daytime (10:00-18:00 local time) and the minimum during nighttime (00:00-06:00 local time) (see also Takaya et al., 2010; While et al., 2017).

The scatter plots of DOISST, SEVIRI and model vs in situ-measured DWA have been computed for the years 2019-2020 (Fig. 6) and organized during spring-summer and winter-autumn seasons (Fig. 7). This choice is aimed at comparing the behaviour of the three products as a function of the seasons, since larger DWA intensities are expected in the spring-summer period.

https://doi.org/10.5194/essd-2021-462
Preprint. Discussion started: 20 January 2022

Overall, there is a good agreement between DOISST and drifter DWAs (Fig. 6a) as confirmed by an almost null mean bias (-0.02 K), low RMSD (0.38 K) and high correlation coefficient (0.82). The largest DW amplitudes reach values as high as 4 K in both DOISST and drifter data. SEVIRI (Fig. 6b) shows the same bias (-0.02 K) of DOISST in reconstructing DWAs but higher RMSD (0.49 K) and lower correlation (0.74). It is relevant to note that the spread of SEVIRI DWAs around the line of perfect agreement is reduced in DOISST, which coherently has a lower RMSD. The model (Fig. 6c) clearly underestimates diurnal amplitudes larger than 1 K, and it is characterized by the highest mean bias (-0.23 K) and RMSD (0.66 K), and lowest correlation coefficient (0.66).

The majority of DWA events lie between 0-1 K all over the year, but higher values are effectively reached during spring and summer (Fig. 7). During these seasons, it appears more evident the capability of DOISST to better describe DWAs larger than 1 K (mean bias = -0.04 K; RMSD = 0.42 K; corr. = 0.83) compared to SEVIRI (mean bias = -0.05 K; RMSD = 0.53 K; corr. = 0.76) and especially to the model (mean bias = -0.27 K; RMSD = 0.65 K; corr. = 0.63). A similar behaviour is obtained during winter and autumn when DWA events exceeding 1 K are also observed, and such intense amplitudes are not found in the model-derived DWAs. Additionally, the good agreement between DOISST and drifters still confirms that interpolated data do not suffer from the increased cloud cover during winter and autumn periods.

(a) (b) (c)

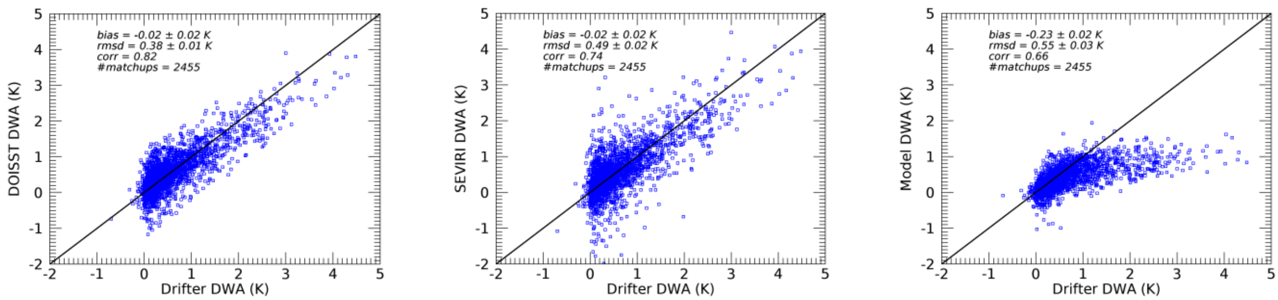


Figure 6. DWA scatter plots for (a) DOISST, (b) SEVIRI L3C and (c) model vs drifters over the period 2019-2020.

https://doi.org/10.5194/essd-2021-462
Preprint. Discussion started: 20 January 2022

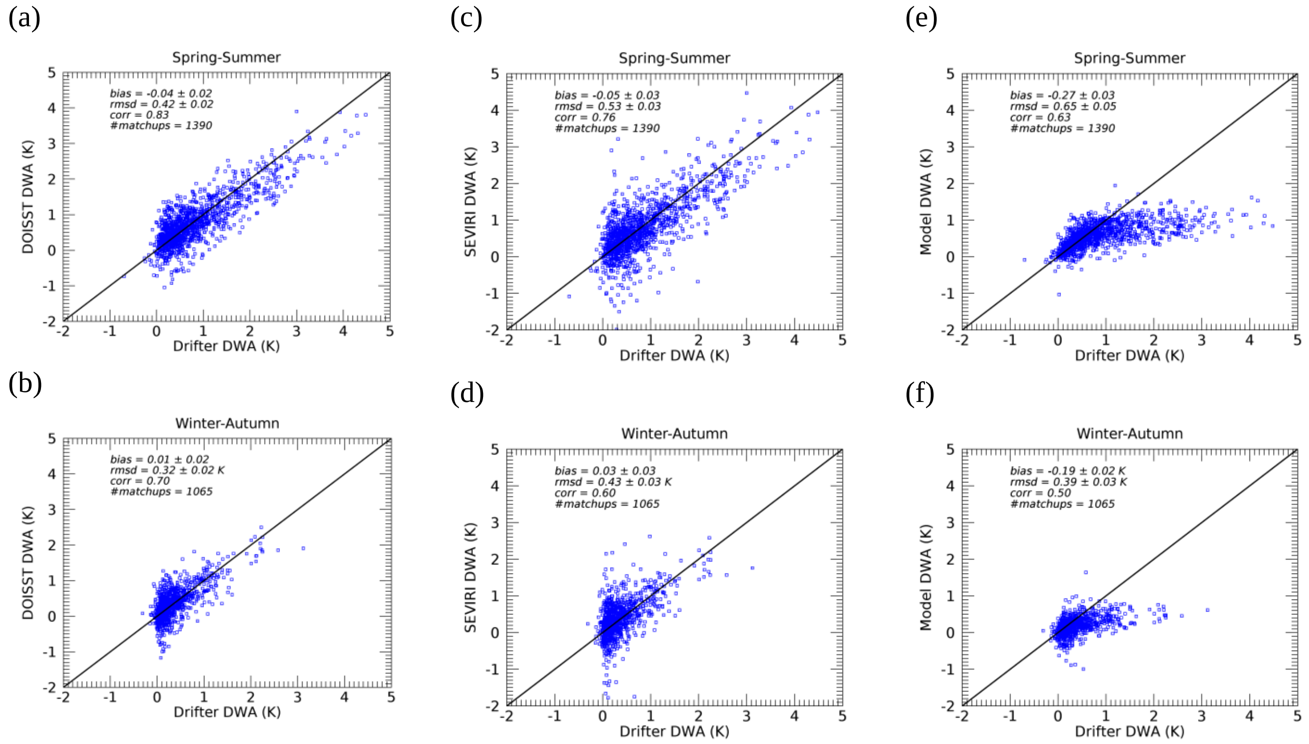


Figure 7. DWA scatter plots for DOISST (a,b), SEVIRI L3C (c,d) and model (e,f) vs drifters during Spring (M-A-M) and Summer (J-J-A), and Winter (D-J-F) - Autumn (S-O-N), over the period 2019-2020.

Having demonstrated the reliability of DOISST in the DWA estimate, we analyze its capability to reproduce the typical spatial variability and intensity of DW events in the Mediterranean Sea, a basin characterized by a frequent occurrence of intense DW events (Böhm et al., 1991; Buongiorno Nardelli et al., 2005; Gentemann et al., 2008; Merchant et al., 2008). In our investigation area, the 2019-2020 mean DWA ranges from a minimum of 0.4 K in the Atlantic ocean box off the Strait of Gibraltar, to a maximum of 1.2 K in several regions of the Mediterranean Sea (Fig. 8a) where individual diurnal warming events exceeding 1 or even more than 2 K are quite frequent (Minnet et al. 2019; Marullo et al. 2016; Marullo et al 2014; see in particular Fig. 3 in Minnet et al., 2019]. The largest DWA were observed in the Levantine Basin, in the North Adriatic Sea and in correspondence with the Alboran Gyre. Less intense, though still remarkable, mean DWA patches reaching 0.9 K are found around the southern tip of the Italian Peninsula as well as in the coastal Ligurian Sea. In the same areas, it is found that the frequency of DW events larger than 1 K and 2 K can reach up to 55% and 10% of the analyzed time series, respectively (bearing in mind that our time series is given by the total number of days in 2019 and 2020) (Fig. 8b-c).

https://doi.org/10.5194/essd-2021-462
Preprint. Discussion started: 20 January 2022

The magnitude of the maximum SST diurnal oscillation is also investigated. The spatial distribution of the maximum DWA observed through 2019-2020 in the Mediterranean Sea (6°W to 36°E and 30°N to 46°N) (Fig. 8d) shows that the largest amplitudes reach and exceed 3 K in 98% of the basin and local DWA patches exceeding 6 K are also ubiquitous, confirming that the Mediterranean is one of the areas with the largest DWs of the global ocean (Minnet et al. 2019, and references therein).

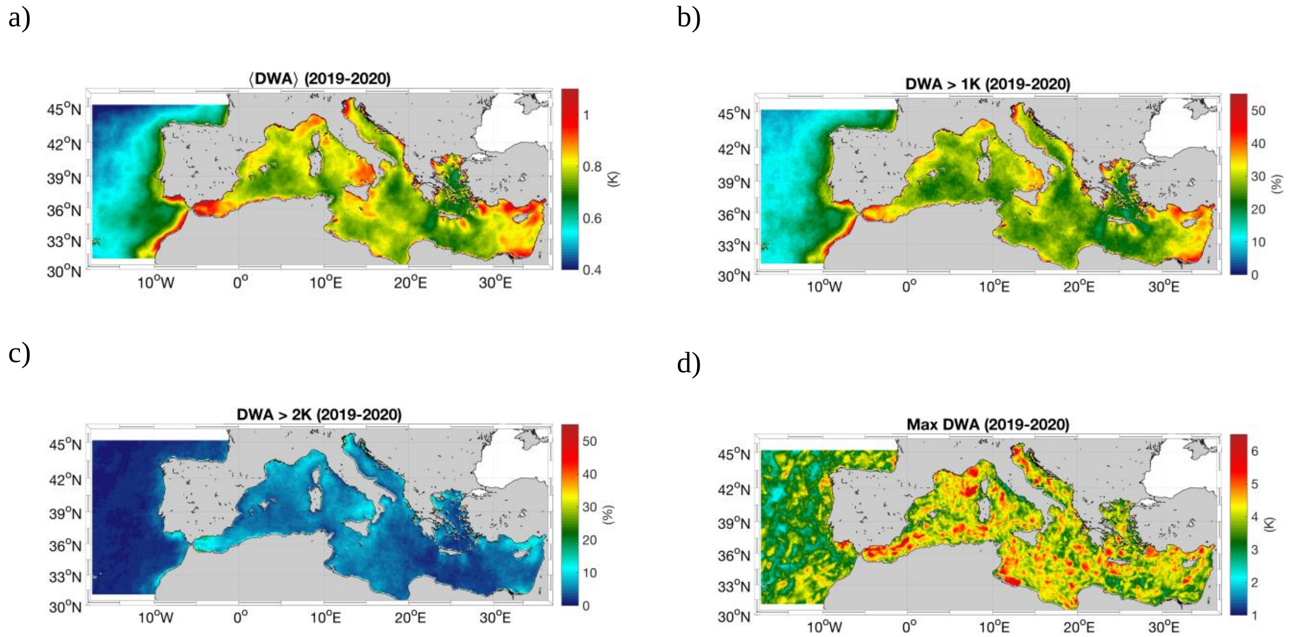


Figure 8. a) Mean diurnal warming amplitude (DWA) derived from DOISST; b) Percentage (over the total number of days in the 2019-2020 period) of DOISST DWA larger than 1 K; c) Percentage of DOISST DWA larger than 2 K; d) Maximum observed DOISST DWA. All the maps refer to the 2019-2020 period.

When compared to the model, DOISST exhibits mean DWAs with larger intensity than model outputs in all the locations of the study area (Fig. 9). The ΔDWA, defined as DWA_{DOISST} minus DWA_{Model}, is always larger than 0.2 K and locally reaches extreme values of ~1 K. The extent of the ΔDWA generally increases in areas where the DOISST mean DWA is larger, such as in the Alboran Sea, Ligurian Sea, Levantine Basin and Southern Tyrrhenian, suggesting a tendency of the model to underestimate the largest DW events.

https://doi.org/10.5194/essd-2021-462
Preprint. Discussion started: 20 January 2022

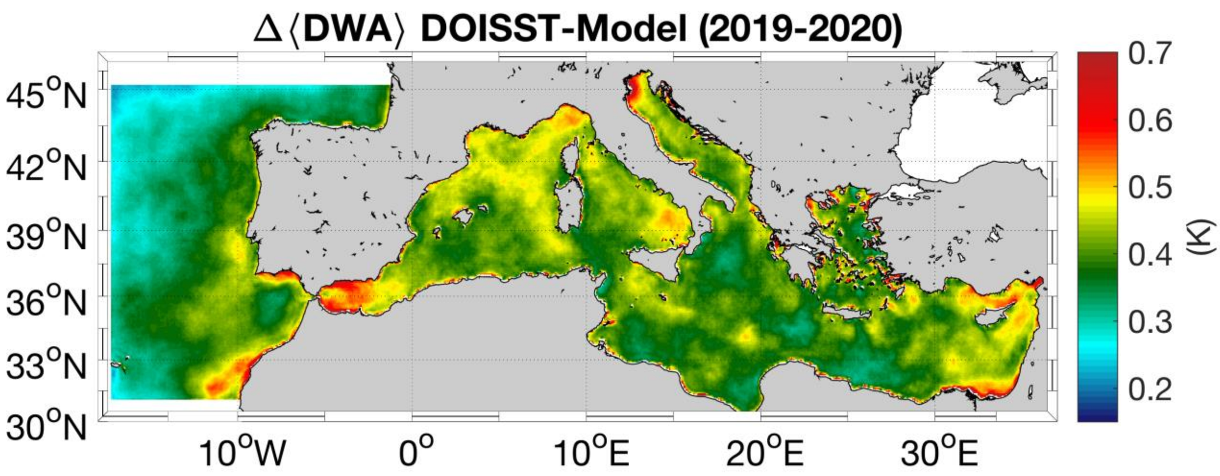


Figure 9. Mean amplitude of the SST DW. Differences between the mean DWA seen by the DOISST product and the model outputs (first layer).

5 Data availability

The Mediterranean diurnal optimal interpolated SST product is distributed as part of the CMEMS catalogue, and identified as SST_MED_PHY_SUBSKIN_L4_NRT_010_036 (CMEMS product reference) and cmems_obs-sst_med_phy-sst_nrt_diurnal-oi-0.0625deg_PT1H-m (CMEMS dataset reference) (https://resources.marine.copernicus.eu/product-detail/SST_MED_PHY_SUBSKIN_L4_NRT_010_036/INFORMATION, last access: 03 November 2021, https://doi.org/10.25423/CMCC/SST_MED_PHY_SUBSKIN_L4_NRT_010_036; Pisano et al, 2021). Access to the product is granted after free registration as a user of CMEMS at https://resources.marine.copernicus.eu/registration-form (last access: 03 November 2021). Once registered, users can download the product through a number of different tools and services, including the web portal Subsetter, Direct-GetFile (DGF) and FTP. A Product User Manual (PUM) and QUality Information Document (QUID) are also available as part of the CMEMS documentation (https://resources.marine.copernicus.eu/product-detail/SST_MED_PHY_SUBSKIN_L4_NRT_010_036/DOCU MENTATION, last access: 03 November 2021). Eventual updates of the product will be reflected in these documents. The basic characteristics of the DOISST product are summarized in Table 1. The reduced subset used here for validation and review purposes is openly available at https://doi.org/10.5281/zenodo.5807729 (Pisano, 2021).

6 Summary and conclusions

A new operational Mediterranean diurnally varying SST product has been released (May 2021) within the Copernicus Marine Environment Monitoring Service (CMEMS). This dataset provides optimally interpolated (L4) hourly mean maps of sub-skin SST over the Mediterranean Sea at 1/16° horizontal resolution, covering the period from 1st January 2019 to near real time (1 day before real time) (Pisano et al., 2021). The diurnal optimal interpolated SST (DOISST) product is obtained from a blending of hourly satellite (SEVIRI) data and model outputs via optimal interpolation, where the former are used as the observation source and the latter as background. This method has been firstly proposed by Marullo et al. (2014), validated over one year (2013) in Marullo et al. (2016), and implemented here operationally. The validation of the operational product was also extended over two years (2019-2020).

DOISST proved to be rather accurate when compared to drifter measurements, and correctly reproduced the diurnal variability in the Mediterranean Sea. The accuracy of DOISST results in an overall, almost null, mean bias of ~0.04 K and RMSD of ~0.41 K (Table 2). This product is also more accurate than the input model, which shows a mean bias of ~-0.1 K and RMSD of ~0.47 K. A warm (positive) and cold (negative) bias characterizes the DOISST and the model, respectively, also during seasons (Fig. 5). These opposite biases are likely related to the different nature of the SST provided by DOISST, model and drifter data, i.e. sub-skin (upper first millimeters from the surface), averaged 1 m depth and 20 cm depth, respectively, and then consistent with the physical consequence of a reduction of the temperature with depth due to the vertical transfer heat process. The DOISST RSMD generally keeps lower values compared to the model, ranging from a minimum of ~0.40 K (vs ~0.42 K for the model) to a maximum of ~0.44 K (vs ~0.56 K for the model).

Compared to its native version (Marullo et al., 2016), the DOISST product maintains the same RMSD (estimated in 0.42 K) but displays a lower mean bias (estimated as -0.10 K). The reduced bias could be ascribed to the fact that valid SEVIRI SST values are always interpolated in DOISST, while they are left unchanged in the original method. The DOISST bias is also lower than that of the OSTIA diurnal product, which produces gap-free hourly mean fields of skin SST for the global ocean, and has been found to underestimate the diurnal range of skin SST by 0.1-0.3 °C (While et al., 2017).

The analysis of the SST diurnal cycle as estimated from both DOISST, model and drifter data shows that the diurnal oscillation in SST is well reconstructed by the DOISST while the model tends to underestimate this amplitude mainly during the central warming hours (Fig. 4), and during spring and summer (Fig. 5). Specifically, DOISST overestimates the mean diurnal amplitude by ~2.3% compared to that of drifters, while the model underestimates it by ~16%. This is particularly evident in the analysis of diurnal warming (DW) events, where diurnal warming amplitudes (DWAs) as estimated by DOISST, model and SEVIRI data are compared vs drifter-derived DWAs. This analysis shows that amplitudes exceeding 1 K as measured by drifters are well reconstructed by DOISST (Fig. 6a) with a mean bias of ~-0.02 K while model SSTs show significantly lower values, with a mean bias of ~-0.23 K (Fig. 6c). The underestimation of the diurnal warming amplitude (DWA) by the model could be related to several factors, such as that the vertical resolution does not resolve the vertical temperature profile within the warm layer, the physics and atmospheric forcing implemented in the model, and/or the assimilation of the foundation SST fields used for the correction of surface heat flux.

The comparison with reconstructed SEVIRI DWAs (Fig. 6b) demonstrates that optimal interpolation does not change the SEVIRI bias, which is practically null for both SEVIRI and DOISST (~-0.02 K), while it reduces the SEVIRI RMSD, from ~0.49 K (SEVIRI) to ~0.38 K (DOISST). The seasonal analysis also reveals that DOISST is not impacted by the different environmental conditions in the Mediterranean Sea, in particular from the much frequent cloudiness during winter and autumn periods.

Overall, the DOISST product is able to accurately reconstruct the SST diurnal cycle, including diurnal warming events, for the Mediterranean Sea and can thus represent a valuable dataset to improve the study of those processes that require sub-daily frequency.

Financial Support

This work has been carried out within the Copernicus Marine Environment Monitoring Service (CMEMS) Sea Surface Temperature Thematic Assembly Centre (SST TAC), contract n° 78-CMEMS-TAC-SST. This contract is funded by Mercator Océan International as part of its delegation agreement with the European Union, represented by the European Commission, to set-up and manage CMEMS.

References

Bernie, D. J., Guilyardi, E., Madec, G., Slingo, J. M., Woolnough, S. J., and Cole, J. (2008). Impact of resolving the diurnal cycle in an ocean–atmosphere GCM. Part 2: A diurnally coupled CGCM. Climate dynamics, 31(7-8), 909-925.

Böhm, E., Marullo, S., and Santoleri, R. (1991). AVHRR visible-IR detection of diurnal warming events in the western Mediterranean Sea. International Journal of Remote Sensing, 12(4), 695-701.

Bowen, M. M., Emery, W. J., Wilkin, J. L., Tildesley, P. C., Barton, I. J., and Knewtson, R. (2002). Extracting multiyear surface currents from sequential thermal imagery using the maximum cross-correlation technique. Journal of Atmospheric and Oceanic Technology, 19(10), 1665-1676.

Bretherton, F. P., Davis, R. E., and Fandry, C. B. (1976). A technique for objective analysis and design of oceanographic experiments applied to MODE-73. In Deep Sea Research and Oceanographic Abstracts (Vol. 23, No. 7, pp. 559-582). Elsevier.

Buongiorno Nardelli, B.; Marullo, S.; Santoleri, R. (2005). Diurnal Variations in AVHRR SST Fields: A Strategy for Removing Warm Layer Effects from Daily Images. Remote Sens. Environ., 95 (1), 47–56. https://doi.org/10.1016/j.rse.2004.12.005.

Buongiorno Nardelli, B., Tronconi, C., Pisano, A., and Santoleri, R. (2013). High and Ultra-High resolution processing of satellite Sea Surface Temperature data over Southern European Seas in the framework of MyOcean project. Remote Sensing of Environment, 129, 1-16.

Chen, S. S., and Houze Jr, R. A. (1997). Diurnal variation and life-cycle of deep convective systems over the tropical Pacific warm pool. Quarterly Journal of the Royal Meteorological Society, 123(538), 357-388.

Clayson, C. A., and Bogdanoff, A. S. (2013). The effect of diurnal sea surface temperature warming on climatological air–sea fluxes. Journal of Climate, 26(8), 2546-2556.

Clementi, E., Oddo, P., Drudi, M., Pinardi, N., Korres, G., and Grandi A. (2017). Coupling hydrodynamic and wave models: first step and sensitivity experiments in the Mediterranean Sea. Ocean Dynamics. doi: https://doi.org/10.1007/s10236-017-1087-7.

Clementi, E., Aydogdu, A., Goglio, A. C., Pistoia, J., Escudier, R., Drudi, M., Grandi, A., Mariani, A., Lyubartsev, V., Lecci, R., Cretí, S., Coppini, G., Masina, S., & Pinardi, N. (2021). Mediterranean Sea Physical Analysis and Forecast (CMEMS MED-Currents, EAS6 system) (Version 1) [Data set]. Copernicus Monitoring Environment Marine Service (CMEMS).

Deser, C., Alexander, M. A., Xie, S. P., and Phillips, A. S. (2010). Sea surface temperature variability: Patterns and mechanisms. Annual review of marine science, 2, 115-143.

Dobricic, S., and Pinardi, N. (2008). An oceanographic three-dimensional variational data assimilation scheme. Ocean modelling, 22(3-4), 89-105.

Efron, B.; Tibshirani, R.J. An Introduction to the Bootstrap; CRC Press: Boca Raton, FL, USA, 1994.

Fiedler, E. K., McLaren, A., Banzon, V., Brasnett, B., Ishizaki, S., Kennedy, J., ... and Donlon, C. (2019). Intercomparison of long-term sea surface temperature analyses using the GHRSST Multi-Product Ensemble (GMPE) system. Remote sensing of environment, 222, 18-33.

https://doi.org/10.5194/essd-2021-462
Preprint. Discussion started: 20 January 2022

Open Access Earth System Science Data Discussions

Gentemann, C. L.,Minnett, P. J., Le Borgne, P., and Merchant, C. J. (2008). Multi-satellite measurements of large diurnal warming events. Geophysical Research Letters, 35, L22602. http://dx.doi.org/10.1029/2008GL035730.

Gentemann, C. L., and Minnett, P. J. (2008). Radiometric measurements of ocean surface thermal variability. Journal of Geophysical Research: Oceans, 113(C8).

Good, S. A., Corlett, G. K., Remedios, J. J., Noyes, E. J., and Llewellyn-Jones, D. T. (2007). The global trend in sea surface temperature from 20 years of advanced very high resolution radiometer data. Journal of climate, 20(7), 1255-1264.

Good, S., Fiedler, E., Mao, C., Martin, M.J., Maycock, A., Reid, R., Roberts-Jones, J., Searle, T., Waters, J., While, J., and Worsfold, M. (2020). The Current Configuration of the OSTIA System for Operational Production of Foundation Sea Surface Temperature and Ice Concentration Analyses. Remote Sensing, 12(720), doi:10.3390/rs12040720

Huang, B., Liu, C., Freeman, E., Graham, G., Smith, T., & Zhang, H. M. (2021). Assessment and Intercomparison of NOAA Daily Optimum Interpolation Sea Surface Temperature (DOISST) Version 2.1. *Journal of Climate*, *34*(18), 7421-7441.

Kotsias, G., & Lolis, C. J. (2018). A study on the total cloud cover variability over the Mediterranean region during the period 1979–2014 with the use of the ERA-Interim database. Theoretical and Applied Climatology, 134(1), 325-336.

Le Traon, P. Y., Reppucci, A., Alvarez Fanjul, E., Aouf, L., Behrens, A., Belmonte, M., ... and Zacharioudaki, A. (2019). From observation to information and users: The Copernicus Marine Service perspective. Frontiers in Marine Science, 6, 234.

Marullo, S., Minnett, P. J., Santoleri, R., and Tonani, M. (2016). The diurnal cycle of sea-surface temperature and estimation of the heat budget of the Mediterranean Sea. Journal of Geophysical Research: Oceans, 121(11), 8351-8367.

Marullo, S., Santoleri, R., Ciani, D., Le Borgne, P., Péré, S., Pinardi, N., Tonani, M., and Nardone, G. (2014). Combining model and geostationary satellite data to reconstruct hourly SST field over the Mediterranean Sea. Remote sensing of environment, 146, 11-23.

Merchant, C. J., Embury, O., Bulgin, C. E., Block, T., Corlett, G. K., Fiedler, E., ... and Donlon, C. (2019). Satellite-based time-series of sea-surface temperature since 1981 for climate applications. Scientific data, 6(1), 1-18.

Merchant, C. J., Filipiak, M. J., Le Borgne, P., Roquet, H., Autret, E., Piollé, J. F., & Lavender, S. (2008). Diurnal warm-layer events in the western Mediterranean and European shelf seas. Geophysical Research Letters, 35(4).

Minnett, P. J., Alvera-Azcárate, A., Chin, T. M., Corlett, G. K., Gentemann, C. L., Karagali, I., ... and Vazquez-Cuervo, J. (2019). Half a century of satellite remote sensing of sea-surface temperature. Remote Sensing of Environment, 233, 111366.

https://doi.org/10.5194/essd-2021-462
Preprint. Discussion started: 20 January 2022

Oddo, P., Bonaduce, A., Pinardi, N., and Guarnieri, A. (2014) Sensitivity of the Mediterranean sea level to atmospheric pressure and free surface elevation numerical formulation in NEMO. Geosci. Model Dev., 7, 3001–3015.

Oddo, P., Adani, M., Pinardi, N., Fratianni, C., Tonani, M., and Pettenuzzo, D. (2009). A Nested Atlantic-Mediterranean Sea General Circulation Model for Operational Forecasting. Ocean Sci. Discuss., 6, 1093-1127.

Oliver, E. C., Benthuysen, J. A., Darmaraki, S., Donat, M. G., Hobday, A. J., Holbrook, N. J., ... and Sen Gupta, A. (2021). Marine heatwaves. *Annual Review of Marine Science*, *13*, 313-342.

Pinardi, N., Allen, I., De Mey, P., Korres, G., Lascaratos, A., Le Traon, P.Y., Maillard, C., Manzella G., and Tziavos, C. (2003). The Mediterranean ocean Forecasting System: first phase of implementation (1998-2001). Ann. Geophys., 21, 1, 3-20.

Pisano, A., Marullo, S., Artale, V., Falcini, F., Yang, C., Leonelli, F. E., ... and Buongiorno Nardelli, B. (2020). New evidence of mediterranean climate change and variability from sea surface temperature observations. Remote Sensing, 12(1), 132.

Pisano, A., Buongiorno Nardelli, B., Marullo, S., Rosalia, S., Tronconi, C., & Ciani, D. (2021). Mediterranean Sea - High Resolution Diurnal Subskin Sea Surface Temperature Analysis (Version 1) [Data set]. Copernicus Marine Environment Monitoring Service (CMEMS). https://doi.org/10.25423/CMCC/SST_MED_PHY_SUBSKIN_L4_NRT_010_036

Pisano, Andrea. (2021). CNR Mediterranean Sea High Resolution Diurnal Subskin Sea Surface Temperature Analysis: Validation subset. https://doi.org/10.5281/zenodo.5807729

Reverdin, G., Boutin, J., Martin, N., Lourenço, A., Bouruet-Aubertot, P., Lavin, A., ... and Lazure, P. (2010). Temperature measurements from surface drifters. Journal of Atmospheric and Oceanic Technology, 27(8), 1403-1409.

Rio, M. H., and Santoleri, R. (2018). Improved global surface currents from the merging of altimetry and sea surface temperature data. Remote sensing of Environment, 216, 770-785.

Storto, A., and Oddo, P. (2019). Optimal assimilation of daytime SST retrievals from SEVIRI in a regional ocean prediction system. Remote Sensing, 11(23), 2776.

Takaya, Y., Bidlot, J. R., Beljaars, A. C., & Janssen, P. A. (2010). Refinements to a prognostic scheme of skin sea surface temperature. Journal of Geophysical Research: Oceans, 115(C6).

Yang, C., Leonelli, F. E., Marullo, S., Artale, V., Beggs, H., Nardelli, B. B., ... and Pisano, A. (2021). Sea Surface Temperature Intercomparison in the Framework of the Copernicus Climate Change Service (C3S). Journal of Climate, 34(13), 5257-5283.

https://doi.org/10.5194/essd-2021-462
Preprint. Discussion started: 20 January 2022

Waters, J., Lea, D. J., Martin, M. J., Mirouze, I., Weaver, A., and While, J. (2015). Implementing a variational data assimilation system in an operational 1/4 degree global ocean model. Quarterly Journal of the Royal Meteorological Society, 141(687), 333-349.

While, J., Mao, C., Martin, M. J., Roberts-Jones, J., Sykes, P. A., Good, S. A., and McLaren, A. J. (2017). An operational analysis system for the global diurnal cycle of sea surface temperature: implementation and validation. Quarterly Journal of the Royal Meteorological Society, 143(705), 1787-1803.