# **Colombian soil texture: Building a spatial ensemble model**

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Abstract. Texture is a fundamental soil property for multiple applications in environmental and earth sciences. Knowing its spatial distribution allows for a better understanding of the response of soil conditions to changes in the environment, such as land use. This paper describes the technical development of Colombia's first texture maps, obtained via a spatial ensemble of national and global digital soil mapping products. This work compiles a new database with 4203-4203 soil

- 5 profiles, which were harmonized at five standard depths (5, 15, 30, 60, and 100 0.5, 5-15, 15-30, 30-60, and 60-100 cm) and standardized with additive log-ratio (ALR) transformation. A stack with compilation of 83 covariates was developed ; including both quantitative and qualitative covariates, and and harmonized at 1 square kilometer of spatial resolution. The top explanatory covariates were selected for each transformation in all standard depths through a recursive feature elimination. Ensemble Machine Learning (EML) algorithms (MACHISPLIN and landmap) were trained to predict the distribution of soil
- 10 particle fractions (SPFparticle-size fractions (PSF) (clay, sand, and silt), and a comparison with SoilGrids (SG) products was performed. Finally, a spatial ensemble function was created , which identified the fewest to identify the smallest prediction errors between EML and SG, then selected the better of these algorithms for each pixel and standard depth. Our results in a digital soil mapping framework signify the first effort to build a national texture map (clay, sand, and silt fractions). The results of EML algorithms show that the accuracy of MACHISPLIN and landmap ensemble algorithms showed that their accuracies
- 15 were very similar in the SPF to the PSF at each standard depth, and both were more accurate than SG. The amount of variance explained (AVE) was between 0.12 and 0.35 for EML, and -0.17 and -0.01 for SG; the concordance correlation coefficient (CCC) was between 0.32 and 0.54 for EML, and 0.04 and 0.16 for SG. The best EML performance was found for the two superficial layers (5 and 15 cm). The accuracy of the largest improvement with the spatial ensemble was higher compared to the other algorithms at all standard depths, but the largest improvement was found at the first layer , where AVE values
- 20 increased between 0.04 and 0.13, and CCC values between 0.04 and 0.10(0.5 cm). EML predictions were frequently selected for each PSF and depth in the total area; however, SG predictions were better when increasing soil depth in some specific regionssuch as Orinoquía and Amazonía. The final error distribution in the study area showed that sand fraction presented higher absolute error values than clay and silt fractions, specifically in eastern Colombia. The spatial distribution of soil texture in Colombia is a potential tool to provide information for water-related water-related applications, ecosystem services, and
- 25 agricultural and crop modeling. However, some aspects must be attended in future efforts to accurately map soil texture; for example, the treatment of future efforts need to improve aspects such as treating abrupt changes in the texture between depths

, unbalanced data, and compositional data consistency in spatial ensemble products and unbalanced data. Our results and the compiled database (??) (Varón-Ramírez and Araujo-Carrillo, 2022; Varón-Ramírez et al., 2022) provide new insights to solve some of the aforementioned issues.

Keywords: Soil Particle Fractions, Ensemble Machine Learning, Compositional Data, Soil Database.

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## 1 Introduction

Soil texture is a complex variable characterized by the amount of claydefined by the proportion of particle-size fractions (PSF), called clay, silt, sand and siltsize particles forming soil aggregates. The spatial variability of soil properties and functions is a key

- 35 component of soil science research, because it allows to understand the response of soil conditions to global environmental or land use changes. Soil textural properties (e.g., particle size soil fraction and distribution) are specifically and sand (Richer-de Forges et al., Soil texture is important to understand soil processes related to agriculture and the environment from the field to the continental scale (Radočaj et al., 2020; Malone et al., 2021; Bönecke et al., 2021; ?)(Radočaj et al., 2020; Malone et al., 2021; Bönecke et al., 2021; C . For example, soil texture is a basic soil variable fundamental soil property for characterizing soil productivity and soil fer-
- 40 tility (Patel et al., 2021; Soropa et al., 2021). Knowledge of soil texture spatial variability is also relevant for assessing soil health and developing strategies to reverse land degradation. Soil texture plays a fundamental role in quantifying the capacity of soils to store carbon and to retain the water required for plants to grow (Dharumarajan and Hegde, 2020; Zhang and Hartemink, 2021). Soil texture is directly measured by soil scientists collecting soil samples and performing soil textural analysis in the field or in the laboratory. One common challenge for soil scientists or pedometricians is to accurately predict
- 45 soil texture (and other soil properties) across areas where no field samples have been collected in the past or at the global scale. At the national to continental scales, there is large uncertainty of current soil texture estimates mainly across large regions and countries of the world where obtaining field soil samples to represent the spatial variability of soil texture is difficult for multiple reasons such as limited funding or site inaccessibility. Additionally, the soil texture study must include two principal statements: this soil property changes with depth increasing (Orton et al., 2016; Poggio and Gimona, 2017) and
- 50 it is compositional data, which means that PSF sum to 100% (%clay+%sand+%silt) and this statement must be satisfied at each location (Amirian-Chakan et al., 2019).

Soil texture estimates are available in different formats such as point data, polygon maps or digital soil maps (Poggio et al., 2021; Anderse --Spatial predictions of soil properties (e.g., elay contentparticle-size fractions proportion) or classes (e.g., soil textural class) across areas where no soil data exists is the main exist is the primary motivation of digital soil mapping (or pedometric mapping)

55 (?)(McBratney et al., 2003). In digital soil mapping, soil texture properties (continuous or categorical) for a specific soil depth and for a given location in the geographical space can be predicted as a function (e.g., empirical function) of the soil forming or weathering environment (climate, organisms, topography, geology, ecology, atmosphere, and human interventions to soils) (Grunwald et al., 2011). The available soil data collected in the field (e.g., content of sand, silt and clay particles) is used to train algorithms or models in combination with gridded environmental information representing the soil weathering environment.

- 60 These layers of environmental information are then used as prediction factors for soil property data (continuous or categorical) extracted from soil samples collected at field. These environmental prediction factors. These environmental prediction factors are commonly acquired from four main\_primary sources: remote sensing, digital terrain analysis, climate, and thematic maps (e.g., soil type, rock type). The use of prediction algorithms or models that are able to can account for the spatial variability of soil distribution is the basis of digital soil mapping , while the spatial support of prediction factors representing the soil forming
- 65 environment is the basis for generating digital soil maps(Wadoux et al., 2021a; Khaledian and Miller, 2020). Digital soil maps are digital soil datasets that represent the continuous nature of soil variability. Predictions of quantitative soil properties (e.g., percentages of clay, silt, and sand) (Liu et al., 2020; Li et al., 2020) or the probability of presence/absence of a soil class (e.g., a soil textural class) (Ramcharan et al., 2018; Kaya and Başayiğit, 2022) are represented on digital soil maps for a given soil depth and for a specific periodof time. These predictions or probability estimates are derived
- 70 probabilities come from the use of statistical models for supervised (at in the presence of training data for a response variable) or unsupervised statistical learning (in the absence of a response variable) (James et al., 2013). Statistical learning methods for supervised learning (e.g., for upscaling soil texture data using digital elevation models) can be applied to categorical (e.g., to solve classification problems) or numerical (to solve prediction problems) datasets (Bischl et al., 2016). Currently, there are literally There are hundreds (if not thousands) of modeling approaches for solving regression and classification problems.
- 75 We could classify these methods in two main into two modeling cultures: one assumes that the data are generated by a given stochastic data model generates the data, and the other uses algorithmic models and treats the data mechanism as unknown (Breiman, 2001). However, it is difficult not easy to classify the large immense diversity of modeling approaches and their possible combinations and therefore. Therefore, the problem of model or algorithm selection to perform predictions in digital soil mapping is an emergent research question.
- 80 Geostatistics and machine learning are the two main-principal forms of statistical learning in digital soil mapping DSM (Hengl and MacMillan, 2019). Geostatistics is a branch of statistics that deals with the values associated with spatial or spatial temporal datasets, whereas spatial-temporal datasets (Webster and Oliver, 2007). In contrast, machine learning is a computer-assisted branch of statistics that uses algorithms developed to solve prediction problems While geostatistics relies on multiple assumptions about the spatial variability of the target variable, machine learning algorithms could be
- 85 considered as assumption-free models. (Witten et al., 2011). Machine learning models are commonly parameterized (selection of multiple modeling parameters) using multiple re-sampling resampling techniques, such as cross-validation or bootstrapping (Brenning, 2012). These resampling techniques allow the algorithm to 'learn' from the data using the capacity of computers to store results from multiple data configurations following the same statistical treatment. This computer-assisted learning allows machine learning algorithms to reproduce the relationship between the response and the
- 90 prediction factors in the statistical space , and can be applied to soil datasets to generate digital soil maps. Geostatistics has historically been the main statistical approach to the prediction problem of soil spatial variability and it is key to determining the spatial structure of error estimates on digital soil maps. In recent decades, machine learning has also become a conventional

approach for digital soil mapping of soil properties and classes. Machine learning algorithms can be roughly divided in into four main groups: a) conventional machine learning based on trees, kernels, linear based or probabilistic algorithms, b) reinforce-

- 95 ment learning algorithms, c) deep learning algorithmsand e, and d) ensemble learning algorithms. The process of extracting information from the use of These ensemble learning algorithms extract information from multiple modeling approaches and combining combine them to create a better solution for a given prediction problem . Machine learning and geostatistics are valuable tools to extract new knowledge around the spatial variability of soil properties (e.g., soil texture) by the means of digital soil mapping. These models and algorithms are rapidly evolving, with most developments or applications being in the
- 100 fields of artificial intelligence, pattern recognition and computer vision. Recent developments on geostatistics and (Yang, 2017)

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Recent developments in ensemble learning efforts have demonstrated great potential for improving the accuracy and spatial detail of current estimates of soil functional properties across scales (Hengl et al. (2021); Wadoux et al. (2020); Llamas et al. (2020)). Several experiences with the mapping of soil texture at different depths have been done: France (Mulder et al., 2016),

105 Scotland (Poggio and Gimona, 2017), Hungary (Laborczi et al., 2019), or China (Liu et al., 2020) are some examples. However, few of them have used spatial ensemble techniques. One representative case was developed by Hengl et al. (2021) for the continent of Africa at three depths (0, 20, and 50 cm) and 30 m spatial resolution.

There is a diversity of emerging research questions in digital soil mapping while coupling the available soil data with the associated prediction factors in the statistical space. Some examples of these questions are: what is the right pixel size?

- 110 (Hengl, 2006) what are the best models and prediction factors? (Guevara et al., 2020). What is the sensitivity of models to the size of training data (Ng et al., 2020), or what are the best big-data management strategies for generating high-spatial resolutionmaps across large areas (e. g., countries) ? (Shangguan et al., 2014) just to mention a few. Here, we compare and test multiple machine learning approaches applied to predict soil texture datasets and provide a solution to identify the lowest prediction bias of multiple soil texture predictions across the national scale of ColombiaThey produced predictions using
- 115 two scale 3D ensemble machine learning (EML) framework; their study utilized an improved predictive mapping framework: spatially-adjusted EML, that better accounts for spatial clustering of points. The spatial cross-validation methodology was a special point of their work, obtaining the following RMSE for  $\approx 122,200$  training samples: clay 9.6%, sand 13.7%, and silt 8.9%. Their results proved to be more accurate than previous works, which is attributable to the addition of higher resolution remote sensing images and Digital Terrain Parameters (DTM), the adoption of methodological improvements in
- 120 hyper-parameter tuning, selection of features, and implementation of ensemble models (Hengl et al., 2021). Models or algorithms for digital soil mapping are evaluated using information criteria and bias indicators that are usually estimated using resampling techniques, such as spatial and nonspatial cross validation (Wadoux et al., 2021b). Some of these accuracy indicators are the root mean squared error, the mean absolute error, the explained variance or the concordance correlation between observed and predicted values, among many others. Here, we use accuracy indicators to evaluate the
- 125 prediction capacity of multiple Colombia has produced maps, either PSF or textural classes, at a national scale through conventional mapping and regional scale using digital soil mapping. On a national scale, these maps use a series of delineations based on qualitative soil characteristics, called cartographic soil units (CSU). This studies have been produced in different

periods: Cortés et al. (1982) (scale 1:5.000.000), IGAC (2003) (scale 1:500.000), and IGAC (2015) (scale 1:500.000). The map carried out by IGAC (2015) represented the PSF through four textural groups of soils (very fine, fine, medium, and

- 130 coarse) in a layer from 0 to 50 cm; however, this methodology ignores the spatial variability inside the CSU. On a regional scale, Araujo-Carrillo et al. (2021) used machine learning algorithms to generate predictions of soil texture across Colombia. Our objective is the development of a digital soil texturedataset across the country. We present a new dataset composed by harmonized measurements of soil texture (n=4203)that we collect from multiple soil surveys. We harmonize the new dataset with gridded environmental prediction factors and compare the accuracy of our predictions against global datasets. Finally, we
- 135 develop a pixel wise show the spatial distribution of clay (%) and its prediction error. However, they ignore the statement of compositional data of the soil texture, and their study just included the surface layer of the soil (0-20cm).

In this work, first, we compared and tested two ensemble machine learning approaches applied to predict soil texture at national scales in Colombia. Second, we compared our results with the global product SoilGrids (SG). Third, we built an ensemble map developing a pixel-wise solution to identify the method with lower prediction bias of multiple soil texture predictions as previously suggested (Gavilán-Acuña et al., 2021).

error. We hypothesized that multiple prediction algorithms are able to could capture the spatial variability of soil texture differently, because they treat the data in different ways to solve prediction problems (e.g., using decision boundaries or probability thresholds or hypothesis of the empirical relationship between the response and the prediction factors). Understanding which are the prediction algorithms and approaches yield lower error levels at the pixel level could benefit model selection efforts in

- 145 digital soil mapping. The digital soil mapping dataset includes: a) point data to calibrate predictions in the form of a regression matrix, b) environmental prediction factors for soil texture and c) digital soil texture maps and their associated uncertainty. Our model predictions improved the accuracy of global available estimates (e.g., SoilGrids250m). The main implication of this study is the opportunity to determine the geographical areas where our results improve the accuracy of previous estimates. The digital soil texture dataset is publicly available at: doi:10.6073/pasta/6dded07af834834ee21a134b247507fd (?). The digital soil
- 150 texture maps across multiple soil depths are also available without restrictions at: doi:10.6073/pasta/91441b598d4480091cd7f86f5d3762bf (?). This work can have positive implications by increasing the quality of, quantity of and access to soil texture information across Colombia.

#### 2 Methodology

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A total of five major stepswere performed by this workOur workflow contains five significant steps: harmonization and transfor-155 mation of soil data, adjustment and selection of covariates, spatial prediction with different algorithms, validation, and spatial ensemble. These sections will be discussed in detail below.

# 2.1 Dataset

A total of 4203-4,203 georeferenced (EPSG:4326) soil profiles were collected from Sistema de Información de Suelos de Latinoamérica y el Caribe - SISLAC, a soil information system developed by FAO (FAO, 2020)-Soil, that all contained information

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160 <u>about soil</u> particle-size fractions (PSF) such as (clay, sand, and siltwere collected, including geographical coordinates (EPGS: 4326). The soil data covered five natural regions (geographic division made based on climatic, vegetation, reliefand soil classes , and soil class conditions) and 31 districts of the continental area of Colombia (Fig. 1) Figure 1) (Rangel-Ch and Aguilar, 1995). The regions wereCaribbean : Caribean in the north, Pacific in the west, Andean in the center (corresponding to the Andes Mountains), Orinoquia in the east, and Amazon in the south(Rangel-Ch and Aguilar, 1995).



Figure 1. Soil-sample points distribution at 5 standard depths in Colombia 0-5 cm depth (Left: training samples, Right: testing samples)

#### 165 2.2 Data harmonization and transformation

Dataset quality was reviewed through consistency inspection. The inspection included two rules: PSF sum equal to ensured by: i) sum of PSF equals 100 and no overlapping between horizons% and ii) no overlapping sampling depth. PSF were transformed at harmonized to five standard depths (5, 15, 30, 60, and 100-0-5, 5-15, 15-30, 30-60, and 60-100 cm), following the vertical discretization as specified in the GlobalSoilMap specifications (?). That transformation wasconstructed (Arrouays et al., 2014)

- 170 The distribution of soil profiles by depth was: 4,203 at 0-5 cm, 4,201 at 5-15 cm, 4,153 at 15-30 cm, 3,974 at 30-60 cm, and 3,597 at 60-100 cm. The soil information for each depth was obtained using a quadratic function of depth with equal areas (spline) (Bishop et al., 1999), through the mpspline function of the agp package (Beaudette et al., 2013) of R version 4.0.3. PSF are compositional data and therefore require special treatment before spatial prediction. PSF at each A D-part composition is an element where all its components are strictly positive real numbers, they stock relative information, and these components
- 175 must sum to 100% (Amirian-Chakan et al., 2019). In this way, soil texture is a 3-part composition, which means that PSF sum to 100% (%clay+%sand+%silt), and this statement must be satisfied at each location. In order to address this statement, PSF at each profile in standard depth were transformed based on additive log-ratio (ALR) transformation (Aitchison, 1986)<del>, instead of</del> other log-ratio transformations such as isometric, centered and symmetry. The properties of the transformations when applied to regionalized compositions were discussed by Pawlowsky-Glahn and Olea (2004). ALR is commonly used for mapping soil PSF
- 180 (Odeh et al., 2003; Poggio and Gimona, 2017; Wang et al., 2020; ?)(Odeh et al., 2003; Poggio and Gimona, 2017; Wang et al., 2020; Li e , preserving information about spatial correlation, showing a distribution more likely to be closer to a normal distribution (Li et al., 2020), and maintaining the compositional aspect of the variables (Lark and Bishop, 2007).

Let  $z_i$ , i=1, 2, 3 (D) represent the clay, sand, and silt fractions, where D = 3 is the number of soil particle-size categories . ALR and inverse ALR transformation are defined as D - 1 = 2 is the number of transformations. ALR transformation is 185 defined in equation 1, and the inverse transformation to obtain the original values of clay, sand, and silt is defined in equation 2:

$$Trans\_i = ln\left(\frac{z_i}{z_D}\right), \ i = 1, 2, \dots, D-1$$

$$\tag{1}$$

$$z_{i} = \begin{cases} \frac{exp (Trans_{i})}{1 + \sum_{j=1}^{D-1} exp (Trans_{j})}, & i = 1, 2, \dots, D-1 \\ \frac{1}{1 + \sum_{j=1}^{D-1} exp (Trans_{j})}, & i = D \end{cases}$$
(2)

where Where  $Trans_i$  is the transformed value of, and  $z_i$  by the ALR transformation. According with Poggio and Gimona (2017) 190 , in this study , clay was used is the original value. According to Poggio and Gimona (2017) and after verifying the possible selection of denominators (normality), this study used clay as the denominator variable. In this way,  $Trans_1 = ln\left(\frac{sand}{clay}\right)$ and  $Trans_2 = ln\left(\frac{silt}{clay}\right)$ . The ALR transformation was implemented using the alr function in Compositional package (Tsagris et al., 2022). The predictive results were back-transformed to PSF the original PSF values (clay, sand, and silt) using alrinv function.

## 195 2.3 Soil covariates

A-Using ArcGIS version 10.3, a total of 83 environmental covariates were selected to broadly reflect soil forming factors, as described by ?: McBratney et al. (2003):

$$S_a = f(s, c, o, r, p, a, n) \tag{3}$$

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where a soil attribute (S<sub>a</sub>) is a function of other properties of the soil at a point (s), the climate (c), organisms (o), relief (r),
parent material (p), age (a), and space (s) (Table 1). The pixel size of the environmental covariates was adjusted to 1 square kilometer using two methods: nearest neighbor and bilinear interpolation. Then, a stack of covariates was created (collection of rasters) was compiled for Colombia.

A recursive feature elimination (RFE) was run for each depth and transformation, using the function rfe of the caret package (Kuhn et al., 2020). The RFE is an algorithm that implements backwards a backward selection of covariates based on

- 205 predictor importance ranking (Kuhn et al., 2020). The goal was to find a subset of covariates that were used to produce the most accurate model possible. A regression matrix for each depth and transformation was built with the selected covariates, and this allowed extraction of the covariate values at the coordinates of each soil sample. With the regression matrix the dataset was divided using a bootstrapping technique (Kuhn et al., 2020) the function createDataPartition of the caret package (Kuhn et al., 2020). This function generates a stratified random split of the data and aims to create balanced splits of the data:
- a part for model training (75 %) (training samples in Figure 1) and another independent part for validation purposes (25 %) (?) (testing samples in Figure 1) (Guevara et al., 2018).

#### 2.4 Prediction models

The spatial distribution of the PSF at each of the five standard depths was modeled through Ensemble Machine Learning (EML) algorithms in two R packages: MACHISPLIN (Brown, 2021) and landmap (Hengl, 2021). EML consists of various approaches based on different methodologies, including stacking methods, averaging methods, bagging, and boosting approaches (Zounemat-Kermani et al., 2021).

MACHISPLIN evaluates different combinations from six algorithms to predict the input data, weighing and evaluating the fit. The interp.rast function of the MACHISPLIN package interpolates noisy multi-variate MACHISPLIN algorithm interpolates multivariate data through EML using six algorithms: boosted regression trees (BRT), neural networks (NN), gener-

- 220 alized additive model (GAM), multivariate adaptive regression splines (MARS), support vector machines (SVM), and random forest (RF). Further, in MACHISPLIN the residuals of the final model are calculated from the full training datasetand these values are interpolated using. This approach evaluates (via k-fold cross validation, where k=10) a method's ability to predict the input data and ensembles of all combinations of the six algorithms weighting each from 0 to 1. The best model will have the lowest Akaike information criterion with a correction for small sizes (AICc). After the best model is determined, the function
- 225 runs the ensemble on the full dataset. Then, residuals are calculated and interpolated using a thin-plate-smoothing splines. This spline; this creates a continuous error surface and is used to correct most of the residual errors in the final ensemble model (Brown, 2021).

#### Symbol SoilFormingfactor Number of covariates

s Soil 28 Soil index (Clay ratio and GSI). Sand and clay mineralogy. Moisture regime. Google Earth Engine, 2020Soil map, seale 1:100.000, IGAC (2012)c Climate 3 1980 – 2011 annual mean precipitation, relative humidity, and average temperature. Climatological database, 1980 – 2011. IDEAM (2015)o Organisms 8 Land cover categories. NDVI, bands 6 and 7 from Landsat 8. Corine Land Cover Classification, seale 1:100.000, 2010–2012. IDEAM (2014)Google Earth Engine, 2020r Relief 23 Digital elevation model (derived parameters). Physiographic landscape and topography. SRTM mission of 2000, at 90 m. Soil map, seale 1:100.000. IGAC (2012)p Parentalmaterial6 Lithology. Soil

#### map, scale 1:100.000.

#### IGAC (2012) a Age 10 Order of soils. Soil map, scale 1:100.000. IGAC (2012) n Space 5 Oblique geographic coordinates. Møller et al. (2020)

No	Factor	Covariate	Source	Scale or resolution	Туре	Unit
1		Clay ratio	Google Earth Engine		Cont	Dimensionless
2	Soil	Grain Size Index - GSI	(2020)	30 m	Cont	Dimensionless
3-22	(s)	Alofana-Kaolinite, Kaolinite, Kaolinite-Alofana, Kaolinite-Gibsite, Kaolinite-Integrated, Kaolinite- Micas, Kaolinite-Montmorillonite, Kaolinite-Quartz, Montmorillonite-Vermiculite, Montmorillonite- Kaolinite, Altered-Quartz, Feldspars-Altered, Feldspars- Amphiboles, Feldspars-Quartz, Feldspars-VolcanicGlass, Quartz-Amphiboles, Quartz-Altered, Quartz-Feldspars, Quartz-Micas, Quartz.	Soil map. IGAC (2015)	1:100,000	Bin	-
23-28		Udic, Ustic, Udic - Aquic, Aquic - Udic, Aquic, Aridic	Soil map. IGAC (2015)	1:100,000	Bin	-
29 30	Climate	Precipitation Relative humidity	Climatological database 1980 - 2011.	1,000 m	Cont Cont	Millimeters Percentage
31	(1)	Mean air temperature	IDEAM (2015)		Cont	Celsius degree
32	Organisms	Pastures, Heterogeneous agriculture, Shrubs, Forests, Permanent crops	Corine Land Cover Classification 2010–2012. IDEAM (2014)	1:100,000	Bin	-
37 38-39		NDVI - index Averages Band 6 and Band 7 - Landsat 8	Google Earth Engine	30 m	Cont Cont	Dimensionless Dimensionless
40-54		Saladares, Mesas, Hills, Glacis, Terraces, Filas - vigas, Slopes, Fans, Crest, Vallecitos, Hogback, Flood Plan, Dunes, Rocky Glacis, Flattening.	Soil map. IGAC (2015)	1:100,000	Bin	-
55 56 57 58 59 60 61 62	Relief (r)	Elevation Aspect Curvature Drainage Distance Slope Terrain Ruggedness Index - TRI Topographic Wetness Index - TWI Valley Depth	SRTM (2000)	90 m	Cont Cont Cont Cont Cont Cont Cont	Meter Degree Dimensionless Meter Degree Dimensionless Dimensionless Meter
63-68	Parental material ( <i>p</i> )	Alluvial Deposits, Colluvial Deposits,Volcanic Ash, Igneous Rocks,9Sedimentary Rocks, Metamorphic Rocks	Soil map. IGAC (2015)	1:100,000	Bin	-
69-78	Age	Andisols, Alfisols, Entisols, Mollisols, Inceptisols,	Soil map. IGAC (2015)	1:100,000	Bin	-

Landmap The landmap algorithm applies the stacking ensemble type. Stacking (sometimes called stacked generalization or committee machine approach) learns in parallel, then fits a meta-model to predict ensemble estimates (Zhang and Ma,

- 230 2012). The "meta-model" is an additional model that basically combines all individual or "base learners" (Hengl, 2021). The train.spLearner function of the landmap package ensembles the following landmap approach ensembles different machine learning algorithms: a fast implementation of RF (ranger), extreme gradient boosting (xgboost), support vector machines (ksvm), neural networks (nnet), and generalized linear models (GLM) (with Lasso or Elastic Net Regularization (Cross Validated Lambda)(evglmnet)). The landmap package extends the functionality of the mlr 'meta-package' (Lang et al., 2019)
- and is based in on super learner. It is a prediction method designed to find the optimal combination of a collection of prediction algorithms, and its framework is built on the theory of cross-validation and allows for a general class of prediction algorithms to be considered for the ensemble (Polley and Van der Laan, 2010).

The two packages share some machine learning algorithms : RF, NN, SVM; but other are different: BRT, GAM, and MARS (in the case of MACHISPLIN ), and xgboost and cvglmnet (in the case of landmap ). There are other differences:

- 240 MACHISPLIN performs k-fold main difference between these two algorithms is the cross-validation(k=10). The best model will have the lowest residual sum of squares during ; MACHISPLIN randomly makes cross-validation, while landmap makes spatial cross-validation(Brown, 2021). While, landmap computes 5-fold cross-validation and then used to determine the meta-learner. Nevertheless, cross-validation is spatial. Randomly splitting spatial data can lead to training points that are neighbors in space with test points. Due to spatial autocorrelation, test. In the random process, the testing and training dataset would not be
- 245 independent in this scenario, with the consequence that cross-validation fails to detect a possible overfitting (Lovelace et al., 2019). That situation is performed by landmap, due it blocks some training points based on spatial proximity to prevent from producing bias dependence (it makes a semivariogram model) to prevent producing biased estimations predictions (Hengl, 2021). However, Wadoux et al. (2021b) indicate that spatial cross-validation methods may provide biased estimates of map accuracy, but standard cross-validation is deficient in the case of clustered data. Additionally, MACHISPLIN constructs the
- 250 best linear model, systematically assigning a weight for each algorithm and evaluating the fit of the ensemble algorithm. In contrast, landmap constructs the meta-model with the predictions of the cross-validation (indicated in the method of ensemble).

## 2.5 Validation

One—The validation process had three steps: i) inclusion of SG layers, ii) calculation of map quality measures, and iii) 255 identification of the spatial distribution of the prediction error.

In the first step, one aim of the work was to compare the spatial prediction of techniques mentioned in section 2.4 against the products from SoilGrids (SG)-SG Version 2.0 (Poggio et al., 2021). SG layers-In this step, SG layers (250 m pixel size) of the PSF at each standard depth were downloadedand resampled to the same spatial resolution as the environmental covariates. After, external validation was performed in order to assess their quality, based on the prediction error (difference between predicted and observed values). The quantitative statisticaneed included the

260 predicted and observed values). The quantitative statisticsused included the .

In the second step, the prediction errors, which is the difference between the predicted value and observed value (Brus et al., 2011), were calculated for each testing sample. After that, were calculated some quantitative statistics: mean error (ME), root mean square error (RMSE), amount of variance explained (AVE), and concordance correlation coefficient (CCC). ME measures bias in the prediction and is defined as the population mean of the prediction errors (Yigini et al., 2018). Values -

- 265 <u>values</u> close to 0 indicate that the predictions are unbiased. RMSE is a measure of prediction accuracy, and a perfect model would have a value  $\approx$  0 (Kempen et al., 2012). The AVE measures the fraction of the overall dispersion of the observed values that is explained by the model, with the model explains, and this measure has an optimal value of 1 (Samuel-Rosa et al., 2015). Finally, the CCC measures the level of agreement of predicted values with observed values (relationship 1:1)(Lawrence and Lin, 1989). These metrics were calculated for EML and SG layers., where 1 is a perfect concordance and
  - 270 <u>0 is no correlation (Lawrence and Lin, 1989).</u>

In the two EML algorithms and comparative dataset (SG)third step, the prediction errors were estimated with the validation data at each depth. The prediction error was calculated at each point as the difference between the observed and the predicted value (Brus et al., 2011). Each prediction error or independent residual was interpolated using ordinary kriging (OK), a widely used geostatistical technique that assumes intrinsic stationarity (Webster and Oliver, 2007). These layers were obtained through

275 the automatic adjustment of the automap package (Hiemstra et al., 2009) and showed the approximated error tend of the estimations of each technique and depth. This process was done for prediction errors of the three approaches (MACHISPLIN, landmap, and SG) in the five standard depths and each PSF. The layers obtained in this step were harmonized according to this study's resolution and extent framework.

#### 2.6 Spatial ensemble

- 280 The spatial ensemble was maps and their ensemble errors were generated for each depth from the results of the spatial modelling and its errors. The spatial ensemble was a function created that identified the best result for each depth and each pixel. The created function performed and PSF (Figure 2). First, with the error prediction maps for each PSF, a mean absolute error (MAE) was calculated for each EML and SG. After, a spatial ensemble function was created to perform a conditional evaluation using the maps of all the techniques prediction maps and their respective errors, subsequently, for MAE; at each pixel, it the function
- 285 identified which model had the least error and selected it. In this way, a final (EML or SG) had the minimum MAE and selected the predictions for clay, sand, and silt of it. At last, an ensemble prediction error map was built with the best result including all the created techniques. The quantitative for each PSF doing a mask by model selected and assigning their respective prediction error (calculated in section 2.5). After, the map quality measures statistics were newly calculated by the final calculated maps.

### **3** Results

290 The main product of this work is a group of maps that contains the PSF predictions for each standard depths. The predictions were made with EML algorithms, and these predictionswere compared with SG products. Additionally, a spatial ensemble (with EML and SG)were constructed for each PSF in five standard depths. Following, the principal process steps This study.



Figure 2. Framework used for generating ensemble maps. PSF means: inputs or outputs for clay, sand, and silt

represents the first effort to provide a national map of soil texture within the framework of digital soil mapping. This work used EML algorithms to improve the accuracy of national soil texture predictions, with a fully independent dataset, concerning the
global product (SG). Also, it provided new insights for assessing the quality and accuracy of global soil texture predictions. The main results are going to be shown in the next following subsections.

## 3.1 Soil texture characterization and dataset

The textural classes for each sample point at 5-five standard depths are shown in Fig. ??. Figure 3. In most USDA textural classes, the dataset has soil samples. SL, L, CL, SCL and C Sandy loam, loam, clay loam, sandy clay loam, and clay were the most frequent textural classes in all standard depths, and SI and SIL textural classes, silty and silty loam textural classes were less frequent. However, the silt-size particles increased in the deepest layers. Also, it is easy to identify that, in the dataset, there

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less frequent. However, the silt-size particles increased in the deepest layers Also, it is easy to identify that, in the dataset, there are not many soil samples with high content of clay and silt, while extreme contents of sand fraction are frequent.

Descriptive statistics are shown in Table 2 for PSF and its transformations (Trans\_1 and Trans\_2). In all textural fractions, the minimum contents were 1% or less for 5 standard depths and the maximum contents for silt fractions were less than elay

- 305 and sand The PSF covers the entire range of measures (0 to 100 %), which is expected for national-scale analysis. The mean and median for sand fraction was higher than clay and silt fractions in all standard depths. However, the differences between the median for clay, sand and silt in the deepest standard depths were less than the most superficial. The SD, indicating that sand is the dominant particle-size fraction in Colombian soils. The standard deviation grows for all fractions in the last two depths and deepest layers; sand content was the fraction with highest SD the highest variation in all depths. Kurtosis coefficients were
- 310 negative for sand fraction in all depths and for the clay fraction in the last three depths; in contrast, for silt fractions, the kurtosis were always positive These results suggest that sand fraction has the highest variability in the PSF for Colombian soils, which rises with increasing depth. The skewness coefficients were positive and less than 1 for all textural fraction datasets., and the sand fraction showed less deviation from the normal distribution, except for 60-100 cm layer. Regarding transformations, the ranges took values from negatives to positives with means around zero; this signifies that ALR transformation improved the
- 315 sand and silt distribution for all depths, except the sand in the first three layers.

## 3.2 Covariate selection

The selection of the best covariate for each transformation was made, in all standard depths, using a recursive feature elimination (rfe) algorithm (Table 3). Ten rfe Ten recursive feature-elimination models were obtained, and individual covariate stacks were built for each transformation in all standard depths. For each standard depthlayer, covariates were selected for Trans\_1 and

- 320 Trans\_2, respectively: 44 and 83 (5-0-5 cm), 54 and 54 (15-5-15 cm), 59 and 83 (30-15-30 cm) 56 and 58 (60-30-60 cm), and 56 and 83 (100-60-100 cm) (Table 3). The Top-top 5 covariates for each selection (Table 3) included soil forming factors: climatic (TEM, RH and PPTtemperature, relative humidity, and precipitation), topographic (altitude, slope, and presence of Flood Planesflood planes), parent material (presence of Alluvial materialsalluvial deposits), activity of organisms (bands 6 and 7 of Landsat 8) and previous soil index information (Clay ratio and SGI). Only Grain Size Index). It is important to highlight
- 325 that the covariates selection had only two binary covariates (Alluvial and Flood Plains)were represented in this selectionalluvial and flood plains).

#### 3.3 Soil texture predictions and SG products validation

Boundary adjustment parameters of external validation Map quality measures are given in Table 4. Referring to ME, the highest values were found in silt fraction, principally in the 100 cm layer with 3.16% and 2.82% for MACHISPLIN and landmap

- 330 algorithms, respectively. Clay and sand fractions for both algorithms showed negative MEvalues (underestimation), except for clay content in the 100 silt fraction was overestimated (positive ME), and clay and sand fractions were underestimated (negative ME); this happened for all depths and EML algorithms, except clay in 60-100 cm layer. Also, clay and sand in the two deepest layers showed the closest to zero ME values For sand fraction, the bias values of landmap predictions were fewer than MACHISPLIN predictions. The ME values closets to zero were found for clay in the deepest layer, suggesting unbiased
- 335 predictions for both MACHISPLIN and landmap For algorithms. Comparing MACHISPLIN and landmap, the RMSE values were similar for each PSF and depth, with the lowest RMSE values found at 15 cm. However, comparing RMSE values between PSF, RMSE values for sand were always higher than clay and silt.



Figure 3. Particle-size soil samples representation in a textural diagram for each standard depth. C: Clay, SC: Sandy clay, SCL: Sandy clay loam, CL: Clay loam, SIC: Silt clay, SICL: Silt clay loam, L: Loam, SIL: Silt loam, <u>SSI</u>: Silt, SL: Sandy <u>Loamloam</u>, LS: Loamy sand, and S: Sand.

The AVE values were from 0.12 to under 0.35. For all standard depths and algorithms, sand fraction had higher AVE values than clay and silt, except in the 60-60-100 cm layer, where AVE for sand and silt were equal. For MACHISPLIN algorithm,
the AVE decreased with increased depth. In contrast, for landmap algorithm, the behavior of AVE values were not correlated with depthyalues for clay were high. In general, for MACHISPLIN and landmap, the capacity of each model to explain the variance decreased when increasing depth; however, it is important to highlight the AVE values for silt fraction in 0-5 and 5-15 cm layers, which were so close to zero. On the other hand, the CCC values were from 0.32 to 0.54. The CCC values for sand were higher than clay and silt for all depths and algorithms, and in all cases. Also, the lowest CCC value between standard depths was found in the deepest layer (100 cm). 60-100 cm), where the data set had fewer sample points than the superficial layers. These results suggest that, for sand, the predicted and observed values agree more than the clay and silt fraction.

Standard	PSF and	Min	Maa	Maaa	Madian	CD	Variation	Cl.
depth (cm)	transformations	WIIN	Max	Mean	Median	50	KULTOSIS	Skewness
	Clay	0.02	95.07	27.88	25.17	17.23	0.43	0.82
	Sand	0.13	99.19	42.69	42.00	22.45	-0.83	0.13
5	Silt	0.10	83.65	29.43	28.00	13.95	0.00	0.46
	Trans_1	-6.06	7.68	0.44	0.50	1.44	1.20	-0.36
	Trans_2	-6.33	7.87	0.14	0.15	0.89	4.19	-0.22
	Clay	0.44	94.50	28.64	26.07	17.07	0.28	0.76
	Sand	0.12	98.00	42.02	41.53	22.18	-0.83	0.17
15	Silt	1.00	81.74	29.33	28.00	13.61	0.00	0.47
	Trans_1	-6.42	4.63	0.38	0.43	1.40	1.01	-0.40
	Trans_2	-3.45	0.11	0.12	4.82	0.82	1.51	-0.26
	Clay	0.29	94.78	30.47	28.40	17.75	-0.15	0.58
	Sand	0.04	98.00	40.48	38.53	22.62	-0.82	0.27
30	Silt	0.36	76.76	29.05	27.34	13.75	0.05	0.55
	Trans_1	-6.83	5.83	0.27	0.27	1.45	1.05	-0.32
	Trans_2	-3.84	3.33	0.04	0.05	0.83	1.00	-0.07
	Clay	0.01	94.50	32.32	30.10	19.16	-0.47	0.45
	Sand	0.03	99.86	39.10	36.08	23.67	-0.79	0.38
60	Silt	0.05	90.13	28.58	26.13	14.61	0.15	0.66
	Trans_1	-7.47	9.14	0.16	0.14	1.60	1.53	-0.20
	Trans_2	-4.67	6.78	-0.03	-0.05	0.93	1.84	0.21
	Clay	0.06	97.07	32.76	30.71	19.70	-0.55	0.41
	Sand	0.01	99.80	38.57	34.76	24.38	-0.72	0.46
100	Silt	0.14	87.50	28.67	26.24	15.40	0.33	0.71
	Trans_1	-8.43	7.35	0.13	0.08	1.64	1.22	-0.16
	Trans 2	-4.25	4.19	-0.05	-0.07	0.94	1.04	0.14

Table 2. Descriptive statistics of PSF and its transformations for each standard depth. Min: minimum; Max: maximum; SD: standard deviation

An evaluation of SG (250 m) products with the dataset validation (the same used for EML validation) was made, and is also is shown in Table 4. Negative MEvalues (underestimation) were found with a range between -8.28% to -6.21% About ME, the highest bias values were found for sand fractionand positive values with range between 0.89% to 3.42% and 3.47% to 4.78% for siltand elay, respectively. This, followed by clay and silt. For sand fraction, the estimations were underestimated (negative ME) in all standard depths; in contrast, for clay and silt fractions, the estimations were overestimated (positive ME), except by silt in 30 - 60 cm layer. These ME values were higher than EML predictions. In respect to Concerning RMSE values, the sand fraction had higher values than the clay and silt fractions fraction, and the RMSE increased with increased depth, which

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Standard depth (cm)	Variable Covs selected		RMSE	Top 5 Covs selected		
=	Trans_1	44	1.21	TEM, RH, PPT, Altitude, L8 b7		
3	Trans_2	83	0.76	TEM, RH, PPT, Altitude, Alluvial		
15	Trans_1	54	1.17	TEM, Altitude, RH, PPT, Clay ratio		
15	Trans_2	54	0.74	TEM, RH, PPT, GSI, Altitude		
20	Trans_1	59	1.24	TEM, RH, Altitude, Flood plane, PPT		
30	Trans_2	83	0.73	Alluvial, PPT, RH, TEM, Altitude		
(0	Trans_1	56	1.34	RH, TEM, PPT, Altitude, Slope		
00	Trans_2	58	0.82	L8 b6, PPT, Alluvial, RH, Clay ratio		
100	Trans_1	56	1.34	PPT, RH, TEM, Altitude, L8 b6		
100	Trans_2	83	0.84	L8 b6, PPT, GSI, Alluvial, Clay ratio		

 Table 3. Top 5 covariates selection for each transformation and standard depth. TEM: temperature, RH: relative humidity, PPT: precipitation,

 L8 b7: Landsat 8 band 7, L8 b6: Landsat 8 band 6, GSI: Grain Size Index.

was equal in EML predictions increasing depth. For AVE, negative and close to zero values were found in all standard depths
 and fractions (-0.17 to -0.01), which were fewer than EML results-0.27 to -0.08). Similarly, the CCC values were close to zero (0.04 to 0.16), and the highest values were obtained for sand and silt fraction in the three most superficial layers.

In Figure 4, there is a visual comparison of our results and SG products. In this representation, we can see some aspects: first, the ranges of the predicted values are wider for MACHISPLIN and landmap than SG (low contrast of colors). Second, the general pattern of PSF distribution is different between our results and SG products. Principally, in the Andean and Caribean

360 regions, our results suggest that these areas have higher values of sand (blue colors) than clay and silt fractions, while SG suggests soils with more content of clay and silt (orange and green colors); also, in Orinoquia region, our results show soils with high content of silt (green colors) and SG displays soils with more content of clay and sand (purple colors). Third, in the southern and eastern areas, the SG products do not have artifacts generated for the prediction of the algorithms; this is an advantage of SG products.

#### 365 3.4 Spatial ensemble

The Fig. 4, 5, 6, 7, Figure 4 to 8 display the final spatial ensemble maps for each PSF, which contain their final error and the model selected at for each pixel at each standard depth. The spatial ensemble, which, as described above, is a collection of best-fit data from 3 separate different algorithms (MACHISPLIN, landmap, and SG), contained common elements/features in most standard depths.

370 The MACHISPLIN algorithm (Yellow colors) was the model with less percentages errors of clay at 5, 15, 30 and 100 cm ; at 60 cm, this algorithm was the best fit for both sand and clay. MACHISPLIN had representation in all natural regions, and In all standard depths, predictions of MACHISPLIN (Yellow color) and landmap (Green color) represented the PSF distribution more than SG (Gray color). The SG selection increased with depth, especially in 60-100 cm layer. The most significant areas where SG had the fewest prediction errors were in Orinoquia and Amazon regions. Similar to SG, the landmap predictions

375 were frequently selected in the deepest layers . These overall in the southern areas of Colombia. MACHISPLIN predictions were most commonly chosen in areas that were extensive and continuous areas, such as Orinoquía and Amazonía region for sand at 60 cm and Caribe region for clay at 100 cm. At 30 cm layer, MACHISPLIS was the poor predictor for silt (less yellow areas). Amazon regions.

By the other hand, landmap (green colors) was the algorithm with less percentages errors of sand and silt at all standard depths, except by the 60 cm, where this algorithm was the best predictor of clay. The landmap had representation in all natural regions. for silt at the deepest layers, landmap had areasthat were large and continuous; in contrast, for sand, landmap was selected mostly in Orinoquia and Amazonía region. For clay at 5 cm, landmap was the less model selectedConcerning PSF distribution, sand fraction had the highest variation in the geographical space in all depths. The highest mountains in the central areas, and the Sierra Nevada in the areas were not continuous.

385 Concerning to SG (gray colors), with depth raising the SG selection increased, specially at 30 northern, and 100 cm layers. The largest areas that were represented by SG products are mainly in Orinoquía and Amazonía region for sandand silt; In contrast, for clay, SG was chosen in continual areas in the country. This product had the fewest contribution in the most superficial layers. hills in the eastern areas had a significant content of sand; in contrast, the finest fractions (clay and silt) were found principally in valley landscapes between mountain chains in central areas of Colombia and hill landscape in the eastern areas.

The external validation showed an improvement in their metrics vs. the use of using a single algorithm (Table 5). The ME values were closer to zero, showing an improvement in the prediction; however, in this ensemble model, predictions of silt fraction had the highest bias, which is a different behavior of EM, where sand fraction had the most biased predictions. RMSE values decreased for all PSF and standard depths, except for MACHISPLIN at 15 cm in silt fraction. By which means a raising in the precision of the map. On the other hand, AVE values increased with the spatial ensemble model; interesting that the AVE

in the precision of the map. On the other hand, AVE values increased with the spatial ensemble model; interesting that the AVE values for silt fraction at 0 - 5 and 5 - 15 cm were higher compared to landmap predictions (Table 4); also the sand fraction had the most improvement betterment respect to clay and silt fraction. Concern to CCC values, these still equal are still equal to or higher than individual algorithms, except by clay at 5 and 60 cm, and silt at 15 and 100 cm in MACHISPLIN. Finally, concern to ME, these values were fewer than ME values of SG algorithm, however for MACHISPLIN and landmap the behavior of ME increased and decreased randomly with the depth increase.

#### 4 Discussion

In this paper, we developed a new digital soil texture dataset that contains legacy soil data, environmental covariates, and the first digital soil texture maps across Colombia. Soil texture is a key property required for many applications in environmental sciences. Colombia's literature on machine learning applied to soil texture mapping is limited. Our results are based on state of the art accuracy and detail spatial resolution of previous conventional

405 of the art ensemble learning. We improve We improved the accuracy and detail spatial resolution of previous conventional

Depht	Method Method	MACHISPLIN				Soil			
(cm)	PSF-Size Fraction	Clay	Sand	Silt	Clay	Sand	Silt	Clay	Sa
0 - 5	ME	<del>-1.00</del> - <u>0.96</u>	<del>-0.71</del> - <u>0.72</u>	<del>1.71</del> - <u>1.68</u>	-1.88	-0.06	1.94	<del>3.68</del> -2.62	-7.82
	RMSE	<del>15.13</del> - <u>15.00</u>	<del>18.61</del> - <u>18.57</u>	<del>12.07-</del> 11.96	<del>15.03-15.38</del>	<del>18.53</del> - <u>19.09</u>	<del>12.11</del> - <u>13.94</u>	<del>18.24-<u>18.48</u></del>	<del>23.84</del>
	AVE	<del>0.26</del> -0.28	<del>0.34</del> -0.35	<del>0.26</del> 0.28	<del>0.27-</del> 0.24	<del>0.35-0.31</del>	<del>0.26</del> 0.02	<del>-0.07</del> - <u>0.10</u>	<del>-0.07</del> -
	CCC	0.42	0.51	<del>0.44</del> -0.43	0.40	0.54	0.36	0.08	<del>0.13</del>
	ME	<del>-0.30</del> - <u>0.34</u>	-0.75- <u>0.71</u>	1.05	-0.90	-0.47	1.37	<del>3.47-</del> 2.16	-7.15
5 15	RMSE	<del>14.36</del> - <u>14.24</u>	<del>18.05</del> - <u>17.82</u>	<del>11.39-</del> 11.30	<del>14.49</del> - <u>14.77</u>	<del>18.47</del> - <u>19.01</u>	<del>11.70</del> - <u>13.62</u>	<del>17.89</del> - <u>18.41</u>	<del>22.94</del>
5 - 15	AVE	<del>0.31-0.32</del>	<del>0.34</del> -0.36	<del>0.29</del> 0.30	<del>0.29</del> 0.27	<del>0.31</del> -0.27	<del>0.25_0.01</del>	<del>-0.08</del> - <u>0.14</u>	<del>-0.06</del> -
	ССС	0.46	<del>0.52-</del> 0.53	0.46	0.43	0.54	0.36	<del>0.08-</del> 0.06	0.1
15 - 30	ME	<del>-0.50</del> - <u>0.48</u>	<del>-1.09</del> - <u>1.03</u>	<del>1.59</del> -1.50	-0.76	-0.47	1.23	<del>3.62</del> -2.50	<del>-6.95</del> -
	RMSE	<del>15.73</del> - <u>15.75</u>	<del>19.06</del> - <u>19.02</u>	<del>11.81</del> - <u>11.76</u>	15.83	18.93	11.75	<del>18.42</del> - <u>18.64</u>	<del>23.59</del>
	AVE	0.23	<del>0.31</del> -0.32	0.29	0.22	0.32	0.29	<del>-0.06</del> - <u>0.08</u>	-0.05-
	CCC	0.37	0.48	<del>0.45</del> 0.44	0.38	0.51	0.46	0.09	<del>0.13</del>
30 - 60	ME	<del>-0.20</del> - <u>0.04</u>	-0.78	<del>0.98</del> 0.82	-0.08	-0.76	0.84	4.64-3.55	-6.21
	RMSE	17.11	<del>20.74-</del> 20.68	<del>13.11</del> - <u>13.13</u>	17.08	20.87	13.30	<del>20.35</del> - <u>20.54</u>	<del>24.55 (</del>
	AVE	0.23	<del>0.25</del> 0.26	0.25	0.23	0.24	0.23	<del>-0.09</del> - <u>0.11</u>	-0.05
	CCC	<del>0.37-</del> 0.36	0.42	<del>0.41-0.39</del>	0.38	0.43	0.39	<del>0.06-</del> 0.07	0.0
(0, 100	ME	-0.03-0.01	<del>-3.12</del> - <u>3.16</u>	<del>3.16</del> <u>3.15</u>	0.05	-2.87	2.82	4.78-3.17	-8.28
	RMSE	<del>17.35</del> - <u>17.29</u>	<del>22.00</del> - <u>21.89</u>	<del>14.18-</del> 14.09	17.53	22.36	14.24	<del>20.77-</del> 21.44	<del>26.55 (</del>
00 - 100	AVE	<del>0.23-</del> 0.24	0.20	<del>0.13</del> 0.14	0.21	0.17	0.12	<del>-0.10</del> -0.17	<del>-0.08</del> -
	CCC	<del>0.35-</del> 0.36	0.38	<del>0.32</del> 0.31	0.37	0.38	0.34	<del>0.06-</del> 0.07	<del>0.04</del>

 Table 4. Map quality measures of each algorithm for PSF in five standard depths. ME: Mean Error; RMSE: Root Square Mean Error; AVE:

Amount of variance Explained; CCC: Concordance Correlation Coefficient. These map quality measures are based on the validation dataset.

Depth (cm)		0 - 5			5 - 15			15 - 30	
PSF Size Fraction	<del>Clay Clay</del>	Sand-Sand	<del>Silt-Sil</del> t	<del>Clay Clay</del>	Sand-Sand	<del>Silt-Sil</del> t	<del>Clay Clay</del>	Sand Sand	<del>Sil</del>
ME ME	<del>2.54</del> -0.92	<del>-1.02</del> - <u>0.98</u>	<del>1.68</del> -1.81	<del>-0.36</del> - <u>0.31</u>	<del>-0.70</del> - <u>0.61</u>	<del>1.02-0.89</del>	<del>-0.16</del> - <u>0.44</u>	<del>-1.11</del> - <u>0.85</u>	<del>0.96</del>
RMSE RMSE	<del>13.99</del> - <u>14.30</u>	<del>16.91</del> - <u>17.14</u>	<del>11.79</del> - <u>12.61</u>	<del>14.14-<u>13.87</u></del>	<del>16.66</del> - <u>16.94</u>	<del>11.71-</del> <u>11.45</u>	<del>15.68</del> - <u>15.54</u>	<del>18.09</del> - <u>18.55</u>	<del>11.40</del>
AVE AVE	<del>0.37-</del> 0.34	<del>0.46 0.44</del>	<del>0.30-0.19</del>	<del>0.33-</del> 0.35	<del>0.44-</del> 0.42	0.28	0.25	<del>0.23 0.35</del>	<del>0.38</del>
$\underbrace{\mathbf{CCC}}_{\mathbf{CCC}}$	<del>0.19</del> 0.47	<del>0.61-</del> 0.60	<del>0.48-0.42</del>	<del>0.46 0.49</del>	<del>0.59</del> 0.58	<del>0.43-</del> 0.45	<del>0.38</del> -0.39	<del>0.53-0.51</del>	<del>0.48</del>
Table 5. Summary of map quality measures for spatial ensemble model for PSF in five standard depths. ME: Mean Error: RMSE: Root Mean									

Square Error; AVE:Amount of variance Explained; CCC: Concordance Correlation Coefficient. These map quality measures are based on

the validation dataset



**Figure 4.** Color composite map of soil texture fractions predictions with the algorithm selected at two depths: landmap for 5-0-5 and 30-5-15 cm; MACHISPLIN for 15, 60 and 100 cm



Texture fractions at 0-5 cm

Figure 5. Ensemble model, error percentage-error distribution, and best model selected at 15-0-5 cm

## Texture fractions at 5-15 cm



Figure 6. Ensemble model, error-percentage-error distribution, and best model selected at 30-5-15 cm

## Texture fractions at 15-30 cm



Figure 7. Ensemble model, error percentage-error distribution, and best model selected at 60-15-30 cm

# Texture fractions at 30-60 cm



Figure 8. Ensemble model, error-percentage-error distribution, and best model selected at 100-30-60 cm

# Texture fractions at 60-100 cm



Figure 9. Ensemble model, percentage-error distribution, and best model selected at 60-100 cm

previous maps. While many studies focus on mapping soil properties such as pH and organic matter, less fewer studies focus on comparing and testing approaches for maximizing accuracy. We improve the accuracy of national soil texture predictions, with a fully independent dataset, respect to global products. Also, we provide new insights for assessing the quality and accuracy of global soil texture predictions. However, we also identify areas where global predictions suggest lower prediction

- 410 variance. global approaches, such as SoilGrids, for maximizing accuracy. Our results contribute with a nation-wide to a national benchmark of the reliability of global predictions compared to national predictions. We base our soil texture predictions in a soil texture dataset with a transformation process. We deliver this soil texture dataset and demonstrate its applicability using digital soil mapping to describe the geography of Colombia's soil texture. We first discuss the general geography of soil texture across the country and then we compare and discuss our findings with previous work.
- 415 Colombia has a great diversity of soil, which changes soils, and their properties change with depth. In the 5-five standard depths, soil texture in Colombia has representation in all textural classes defined by Staff (2014). As depth increases, the soil thinstexture is finer, and the proportion of clay and silt rises. This is evident in southern, southeastern, eastern, and northeastern Colombia, where Fig. ?? demarcates this with redder colors at 60 and 100 cm. On the other hand, coarse soils (blue colors) are in central and northern areas, and these soil textures hold with increasing depth. This high diversity of soil texture is due to
- 420 the high number of interactions between soil forming factors, specifically particularly the great diversity of parent materials, within Colombia (IGAC, 2015; Araujo et al., 2017).

Some topography and parent material covariates were the principal drivers in texture modeling. The key focal areas with fine and medium textures are found in the northwest (Floodplain and land depressions), in central areas (Magdalena River valley), in the west (Cauca River valley), in the south (Amazon region), and in the east (Orinoquia region). All these regions have **a** 

- 425 common soil forming factor specific soil forming factors, such as alluvial parent material that is deposited by one or many rivers, that are fine soil which are soil fine-size fractions driver (Flórez, 2003). 2003). On the other hand, medium to and coarse textures are principally found in the hillsides of mountain landscapes in central, southern, and southwestern regions. Mainly, these coarse soil textures are due to the presence of sandstone, conglomerate sandstone, granites, and gneisses, among others, that have siliceous and quartz rocks (Catoni et al., 2016), volcanic materials, and glacial clast (IGAC, 2015) that are presented
- 430 in these areas. Despite the relationship between soil texture distribution<del>and relief, relief,</del> and parent material covariates, only altitude (quantitative), slope (quantitative), alluvial (binary)<del>and Floodplane, and flood plane</del> (binary) covariates were <del>shown</del> present in the top 5 predictors for each standard depth.

Although parental material is very important critical in the soil texture spatial distribution, the covariates selection identified that the climatic covariates were more important (i.e., TMED, RH, and PPT). The covariates used to describe the parental

- 435 material were binary class variables, maybe : maybe the following exercises should include quantitative variables to identify this soil forming factor, for example, using radar remote sensing (Niang et al., 2014) or based in on the spectral response in the visible and near infrared near-infrared spectrum (Vis-NIR), medium infrared (MIR), and Vis-NIR-MIR (Campbell et al., 2019). In the PSF predictions (in specific, the ALR components), the importance of the climatic covariates did not have obvious apparent changes with depth. The elimate conditions of the country country's climate conditions have led to relatively strong
- 440 physical weathering in the soil forming process (Osman, 2013). Due to the country<sup>2</sup>'s location, it is influenced climatologically

by the atmospheric circulation of the Caribbean Sea, Pacific Ocean, the Amazon basin, and the orographic barrier of the three branches of the Andes Mountain (Poveda, 2004). Furthermore, in this study, the variables were chosen to maximize the predictive power of the models, not their explanatory capabilities.

- Colombia has not produced maps on PSF at a national scale with DSM productsbut is has developed soils, but Colombia has
  445 developed soil surveys through conventional mapping . These maps use a series of delineated polygons based on qualitative soil characteristics called cartographic soil units (CSU), and have been produced in different periods: Cortés et al. (1982) (scale 1:5.000.000), IGAC (2003) (scale 1:500.000), and IGAC (2015) (scale 1:500.000). The map carried out by IGAC (2015) represented the PSF through four textural groups of soils: very fine, fine, medium, and coarse. In that study, each CSU's group texture was calculated whit a weighted average according with each soil profile representation in the CSU and the
- 450 textural group for each profile was calculated with a weighted average soils fraction between 0 and 50 cm depth according to horizon thickness. The best example of the subnational scale is the work developed by Araujo-Carrillo et al. (2021). In the Cundiboyacense high plateau, their study represented the clay fraction and textural classes only for the surface layer of the soil (0-20 cm) and used the random forest machine learning algorithm.
- The textural soil (IGAC, 2015); then, the textural soil distribution of Colombia presented in this study is not directly comparable with previous national textural soil maps. Due to the methodology used in IGAC (2015), the depth studied is different, the polygons delineated (CSU) have a unique value for an entire area, and CSUs are not an uncertainty value associated. These last two reasons are the major primary use limitations in traditional soil surveys (Angelini et al., 2016). Despite that, the maps produced by this study and those of the IGAC project , both show 2 major show two significant areas with similar attributes. In the northwest (Caribbean region) and southern (Amazon region), the IGAC study present presents a fine group texture (clay
- between 40 and 60%), and this current result shows that levels of clay percentages in that clay range. However, there are is a principal region in the western (Orinoquia region), where the two results are very different. The previous result shows these areas with a coarse textural group, and this current result displays low percentages of sand fractions for 15, 30, and 60-5-15, 15-30, and 30-60 cm depths. These differences are due to the low soil sampling density; where there is just one observation, and in this current study, its nearest predictions are driven by soil data.
- 465 In other countries there have been several experiences with the mapping the PSF at different depths using DSM products: France (Mulder et al., 2016), Scotland (Poggio and Gimona, 2017), Hungary (Laborezi et al., 2019), or China (Liu et al., 2020) are some examples. However, few have used the spatial ensemble techniques. One representative case was developed by Hengl et al. (2021) for the continent of Africa at three depths (0, 20 and 50 cm) and at 30 m spatial resolution. They produced predictions using 2 scale 3D EML framework implemented in the mlr R package. Their study utilized an improved predictive
- 470 mapping framework: spatially-adjusted EML, that better accounts for spatial clustering of points. A special point of their work was the spatial cross-validation methodology, obtaining the following RMSE for ≈ 122.200 training samples: clay 9.6%, sand 13.7%, and silt 8.9%. Their results proved to be more accurate that previous works, which is attributable to the addition of higher resolution remote sensing images and Digital Terrain Parameters (DTM), Regarding map quality measures, RMSE had the highest values for the sand fraction. This is the adoption of methodological improvements in hyper-parameter tuning, feature
- 475 selection, and ensembling of models using the Super Learner algorithm (Hengl et al., 2021). Of the above considerations, we

did not work with higher resolution covariates from earth observations data or DTM derivatives, but we implemented several concepts, like spatial cross validation, feature selection, and the use the EML.

Although it is a challenge to compare the results from several studies, especially as the focus of each study is different, only as a reference point, a comparison was developed to obtain statistics with the works from France, Scotland, Hungary,

- 480 and China. At topsoil depth, the CCC obtained by Mulder et al. (2016) was similar in the sand (0.63), but more accurate in the clay (0.53) and the silt (0.61). The process employed regression tree modelling (cubist), a data-mining technique that was not used in same behavior found by other studies that implemented different algorithms (geoestatistics and machine learning) (Poggio and Gimona, 2017; Laborczi et al., 2019; Liu et al., 2020); however, the joint statement in those studies was that, for sand fraction, the EML employed in ranges were wider and the SD was higher compared with clay and silt fraction. The CCC
- 485 values for sand were higher than clay and silt in all cases, and this is the same behavior found by Mulder et al. (2016) in France. Additionally, for all PSF, they found that the predictions were less reliable for deepest layers; this is the same statement found in this work. In the case of Scotland, between 0 to 100 cm depth, Their map quality measures are better than ours; however, the average RMSE was 7.10% in the clay, 15.86% in the sand and 11.44% in the silt. All the statistics were better soil samples used in that work were between 28.000 and 3000, decreasing with increasing depth.
- 490 The qualitative evaluation for SG at a global scale showed that coarse-scale patterns are well reproduced (Poggio et al., 2021) . Nevertheless, in a quantitative evaluation with Colombian soils, SG products cannot explain the variance (AVE values negative), and they used the hybrid geostatistical Generalized Additive Models (GAMs), combining GAM with Gaussian simulations. About 9000 sampled locations were available for that model in almost 78.000 km<sup>2</sup> (Poggio and Gimona, 2017) . Laborezi et al. (2019) produced maps of PSF on Hungary through composite regression kriging directly compiled (D) and
- 495 synthetized (S), obtaining RMSE in the clay of 8.91% (D)their predictions are not according (CCC close to zero) with our validation dataset, and their RMSE values are significantly higher than ours. Liu et al. (2020) built a national map of silt in China, and 9.32% (S), and the sand of 16.38% (D) they compared their results with SG products through RMSE values. They found that their RMSE values were higher than the RMSE of SG, and 16.92% (S). The results in Hungary were similar in sand that calculated in this work. In China with spatial resolution of 90 m and with 4579 soil profiles, Liu et al. (2020) obtained
- 500 an RMSE of 9.79% in clay, 18.65% in sand and 14.76% in silt, with statistics very similar at other depths. They constructed random forest models with a limited number of sparse soil profiles sites, i.e., nearly one soil profile site per 2000 km<sup>2</sup> on average in their study area. We obtained better results in sand and silt, but worse in clay at the same depth. in many specific areas, SG did not represent the local behavior of the PSF. In this way, we suggest, for applications that need textural soil information at a national scale, to use our results obtained with individual algorithms (MACHISPLIN and landmap) and the
- 505 ensemble maps. However, it is important to stand out that in some areas of Orinoquia and Amazon region, the SG had the fewest prediction error; these regions have in common that there is a low soil sampling density.

In the case of SG 2.0 the RMSE was 13% for clay, 18% for sand and 13% for silt. The qualitative evaluation showed that coarse scalepatterns are well reproduced (Poggio et al., 2021). The SG products showed good results in silt in the spatial ensembled generated, but in specific depths (30 cm and 100 cm) and regions (zones east and south of the country). In other

510 PSF and depths, the spatial ensembled showed better results with the EML worked.

According to the description above, the results of the spatial ensembles were acceptable in general terms, and this This work could identify the <u>PSF's</u> better models, error trends, and prediction layers<del>of the PSF</del>. However, in many areas, depths, and textural fractions<del>the results did not have good quantitative statistics, the map quality measurements are low; for example, we desire to increase AVE and CCC values.</del> The causes can be many: the relations between some soil properties and landscape

- 515 attributes are nonlinear, complex, or unknown, a concept defined by Minasny and McBratney (2010). Linked to the aforementioned is the distribution of the soil samples. The study had an unbalanced representation and spatial clustering, for example; for example, the central zone (Andean region) was the most represented (bias towards potentially productive areas), while the east and southeast zones were the least represented(Fig. 1), so many predictions were largely controlled by point data, then, some large artifacts (e.g. lines and blocky outputs) are shown in these areas, a similar case to that reported by Hengl et al.
- 520 (2014). It is important to say that in the SISLAC database some legacy soil profiles had significant positional errors, due to the proposals of traditional soil surveys carried out by the IGAC.

Overall, this study used the best available environmental information, and it represented the PSF for all country at five standard depths. However, the differential factor included maps that represent the best model (EML or SG) in each area of the country at different depths, called in this work spatial ensembled. Nevertheless, the approach has limitations, for example

- 525 with the results of the spatial ensembled it is not possible to sum 100 with all PSF, because this technique is a combination the best EML or SG worked. Also, there is an abrupt change in the PSF through depths in some areas, a normal behavior in some kinds of soils (for example from alluvial parent material), but not on soils like those found in the peneplain landscapeThese artifacts are derived mainly from covariates related to satellite images, such as bands 6 and 7 from Landsat 8, and its derivatives (Clay ratio and Grain Size Index). These four covariates were obtained from Google Earth Engine with seam carving and had a
- 530 significant importance in the recursive feature elimination model and their score importance (overall in rfe model) are between 10 and 14% respect to the best subset of covariates (Appendix A1). Additionally, for agricultural studies, the use of this results will be straitened to the agricultural limit defined in Colombia, places where the results do not have artifacts (Appendix A2).

In Colombia, DSM has new and great challenges to attend map-user's requirements, such as soil texture predictions with uncertainty improvements and soil maps with better spatial resolution. There are two three principal strategies to improve predic-

535 tions: treatment of unbalanced soil-data soil data, management of PSF transformations, and incorporation of new environmental-covariates environmental covariates related to soil texture drivers. Attending the first strategies, strategy is necessary to raise the soil data base whit database with available soil information from other sources such as detailed soil surveys, soil degradation, and soil management studies made by national and governmental institutions (e.g. IGAC, IDEAM, or UPRA); or obtaining soil textural fractions from other kind the amount of each fraction from other kinds of soil analysis, such as Visible Near InfraRed-Short

- 540 of soil minerals (Lagacherie et al., 2020). Also, model-building processes by soil group (Kempen et al., 2009) or homosoil (Angelini et al., 2020; Malone et al., 2016) (Mallavan et al., 2010; Angelini et al., 2020; Malone et al., 2016) have been used to get pedologically-plausible predictions in areas without high soil-sampling density. Finally, taking heed of second strategies there are Other log-ratio transformations could be applied as a second strategy to improve ALR transformation issues. For example, Wang and Shi (2017) indicated that in some datasets, the changes in the denominator selection in Additive log-ratio
- 545 transformation could represent different predictions and decrease the accuracy of the estimates; then, using centered log-ratio

transformation, this issue could be avoided (Amirian-Chakan et al., 2019); also, data sets with zero values must be threatened with symmetry and isometric log-ratio transformation (Li et al., 2020). Finally, as a third strategy, some qualitative and quantitative environmental covariates that could buttress the predictorsstack, such as depth to bedrock, and soil horizons designations and thickness; also, to improve the visual quality of the results, a previous covariate analysis, such as, principal components (Hengl et al., 2014) or a smoothed strategy.

## 5 Conclusions

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We provided the first comparison of the PSF across Colombia between EML models (MACHISPLIN and landmap) and the <u>SG's</u> existing soil texture maps<del>provided by SG</del>. The study shows that the <u>prediction of the spatial distribution of soil texture</u> prediction with national datasets was<del>better an average</del>, on average, 17% better (in terms of RMSE) using EML models than the

555 SG products. Between MACHISPLIN and landmap, there was no better EML model, because the quantitative statistics were very similar. In function of the PSF, the spatial distributions did not exhibit a fraction with better results. While the silt had the lowest RMSE at different depths, the sand had the highest AVE, and the clay had the lowest ME. However, in the case of the depths, at 5, 15 and 30 cm were obtained the better results for all the PSF, while at 60 However, layers of 0-5, 5-15, and 100 cm the worst, which indicates 15-30 cm obtained the best results, which indicate the effectiveness in the depths closest to the soil

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560 surface.
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Another valuable contribution developed in this study was the implementation of the spatial <u>ensembled ensemble</u> of soil texture <u>fraction\_fractions</u> on a national scale and at different depths. This implementation identified the best result for each depth and each pixel. Although the SG products had the worst quantitative statistics<del>overall</del>, in some areas of the country, these products performed well, <del>such as with 30 cm silt mainly</del> in the south. However, with the spatial ensembled<del>was possible to get</del>, the best composition of the models <del>worked</del>was possible.

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The prediction of the spatial distribution of soil texture fraction obtained particle-size fractions can provide soil information for water related water-related applications, ecosystem services, and agricultural and crop modeling. However, the results had limitations, especially with the abrupt change in the texture fraction through the depths in some areas and the handling of the compositional data (sum equal to 100) in the spatial ensembled productssome artifacts in the southern and eastern areas. Treatment of unbalanced soil-data soil data and incorporation of more appropriate environmental-covariates are key

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environmental covariates are crucial to improving accuracy in the future. Dataset are available at: https://portal.edirepository.org/nis/mapbrowse?packageid=edi.746.1. This repository contains the

data set for each standard depth. For each sample point are shown PSF and ALR transformations (

# 6 Code and data availability

575 Dataset are available at: https://portal.edirepository.org/nis/mapbrowse?packageid=edi.746.2. This repository contains the data set for each standard depth. For each sample point are shown PSF and ALR transformations (Trans\_1 and Trans\_2) (Varón-Ramírez and Araujo-Carrillo, 2022).

Textural soil maps are available at: https://portal.edirepository.org/nis/mapbrowse?packageid=edi.972.3. In this repository the users are going to find 9 raster stacks: PSF obtained with landmap and MACHISPLIN algorithms (2 stacks); PSF obtained

580 from SG (1 stack); residual of the PSF predictions for landmap and MACHISPLIN algorithms and SG (3 stacks); and finally PSF predictions obtained through spatial ensemble technique (3 stacks). All stacks contain information at 5 standard depths (Varón-Ramírez et al., 2022)

Rproject scripts to reproduce the spatial ensemble procedure and models validation area are available at: https://github.com/VimiVaron/Textural-maps-Colombia.git

# 585 Appendix A

# A1 Importance of covariates in recursive feature elimination for Trans\_1 and Trans\_2 prediction

Depth (cm)	
0-5	Trans_2 ) (?) Textural soil maps are available at: https://portal.edirepository.org/nis/mapbrowse?packageid=edi.972
5-15	Trans_2 stacks);
15-30	Trans 1 stack); residual of the PSF predictions for landmap and MACHISPLIN algorithms and SG (3 stacks); and finally
30-60	
60-100	

 Table A1. Representation of importance scores for satellite-derived covariates. GSI: Grain Size Index, L8 b7: Landsat 8 band 7, L8 b6:

 Landsat 8 band 6.

**B1** Agricultural Frontier in Colombia



Figure B1. Compositional texture map (Clay, Sand, and Silt), integrated from 0 to 100 cm, in the agricultural frontier in Colombia.

Author contributions. Varón-Ramírez contributed with conceptualization of soil textural-data management methodology, data cleaning, har-

590 monization of the soil dataset, writing of results and discussion, and elaboration of cartography results. Araujo-Carrillo contributed with environmental covariates construction, writing of methodology, results and discussion, and elaboration of cartography results. Guevara contributed with conceptualization of spatial ensemble models strategies and the comparison with global products and writing the introduction and discussion.

Competing interests. The authors declare that they have no conflict of interest.

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