

# World Atlas of late Quaternary Foraminiferal Oxygen and Carbon Isotope Ratios

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**Abstract.** We present a global atlas of downcore foraminiferal oxygen and carbon isotope ratios available at <https://doi.pangaea.de/10.1594/PANGAEA.936747> (Mulitza et al., 2021). The database contains ~~2,108~~2,106 published and previously unpublished stable isotope downcore records with ~~361,949~~362,067 stable isotope values of various planktonic and benthic species of ~~foraminifera~~Foraminifera from 1,265 sediment cores. Age constraints are provided by 6,153 uncalibrated radiocarbon ages from 598 (47%) of the cores. Each stable isotope and radiocarbon series is provided in a separate netCDF file containing fundamental meta data as attributes. The data set can be managed and explored with the free software tool PaleoDataView. The atlas will provide important data for paleoceanographic analyses and compilations, site surveys, or for teaching marine stratigraphy. The database can be updated with new records as they are generated, providing a live ongoing resource into the future.

## 1 Introduction

Stable oxygen and carbon isotope ratios measured on foraminiferal shells are often regarded as the foundation of Marine Geology and Paleoceanography. The importance of these proxies stems from their broad applicability in time and space, their established and efficient analytical methods and their great value for stratigraphy and paleoceanographic reconstructions (see review by Pearson, 2012). Since the pioneering work of (Urey, 1947), millions of foraminiferal isotope measurements have been performed representing time slices from the Middle Jurassic (e.g. Vetoshkina et al., 2014) into the Anthropocene (e.g. McGregor et al., 2007). Foraminiferal isotopes have substantially contributed to the reconstruction and understanding of the global climate evolution since the Early Cretaceous (Cramer et al., 2009) including the validation of the orbital theory of the ice ages (Hays et al., 1976), reconstructions of ice volume (Shackleton and Opdyke, 1973; Waelbroeck et al., 2002) and water mass structure, ocean circulation and carbon cycling (Curry et al., 1988; Duplessy et al., 1988; Boyle and Keigwin, 1987).

Despite their importance for the understanding of the Earth System, foraminiferal isotope data have not been systematically catalogued [globally](#) or stored in a database in a consistent and standardized format. Foraminiferal isotope data are usually available in arbitrary data formats and scattered across different data repositories, which hinders an automated analysis. Harmonized data collections have the advantage that (i) information about data coverage can be immediately accessed and visualized, for example in the planning phase of research projects, (ii) data can be quickly compared for verification/quality control or to separate local signals from global signals and (iii) that customized software can be used to visualize and analyse the data.

Here we present the first global atlas of foraminiferal stable isotope data (with uncalibrated radiocarbon ages where available). The data are stored in netCDF format (Rew and Davis, 1990) and can be directly analysed and visualized with the free software tool PaleoDataView (PDV, Langner and Mulitza, 2019). In PDV, age information for a specific sediment core is linked to any downcore proxy series imported for that core within the same collection. This strategy ensures the long-term maintainability and consistency of the age models across different proxy records (e.g. stable isotope records of different species) in the same collection. The netCDF format also allows the data to be analysed using programming languages such as MATLAB, R, Fortran, C++ and Python.

## 2 Data sources and harmonisation

The database is provided as a collection of [2,1082,106](#) netCDF files with  $\delta^{18}\text{O}$  and  $\delta^{13}\text{C}$  data of species-specific foraminiferal carbonate and 598 netCDF files with raw radiocarbon ages (see references in Table A1). A detailed description of the attributes and variables in the netCDF files is provided in Supplements S1 and S2. About 79% of the files containing stable isotope records were derived from data downloaded either from PANGAEA ([www.pangaea.de](http://www.pangaea.de)) or NOAA's National Centers for Environmental Information (NCEI, [www.ncdc.noaa.gov](http://www.ncdc.noaa.gov)). The remaining 21% of the stable isotope files are based on data

obtained directly from a stable isotope laboratory by one of the co-authors (8%), or have been digitized from tables provided in papers, or paper supplements (6%), or through personal communication (7%). Radiocarbon data are not as frequently archived in public databases as stable isotope data. Only 62% of the files containing radiocarbon data were obtained from NOAA or PANGAEA, whereas data in 32% of the files were copied from tables in papers or paper supplements. 5% of the files are based on data directly obtained from laboratories, while only 1% was obtained through personal communication. The data set also includes 105 previously unpublished species-specific stable isotope downcore records including both  $\delta^{18}\text{O}$  and  $\delta^{13}\text{C}$  values and 45 species-specific isotope downcore records for which either  $\delta^{13}\text{C}$  or  $\delta^{18}\text{O}$  was previously unpublished (Supplement S3). An Excel spreadsheet containing all data sources is available from Zenodo (<https://zenodo.org/record/5552329>).

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To generate the netCDF files, meta data, isotope data and radiocarbon ages (if available) were first assembled in species- and site-specific Excel files in the format required by PDV. The species names were preserved as used in the original publication. If more than one stable isotope record of the same species was available for the same core, we added a suffix (e.g., size class or version) to the species name. The Excel files were ~~than~~ then edited for units (mainly conversion from “cm” to “m” and years to kiloyears) and meta data were added. Unavailable data fields were filled with “NaN”. Finally, the Excel files were converted to netCDF files using the PDV import tool. Stable isotope data and radiocarbon data were saved in separate files to allow the radiocarbon file to link to several proxy records from the same core via the core label. After import, the data were inspected and quality controlled in PDV. Every row of the downcore data fields is associated with a “use-flag” indicating whether the values should be included in an analysis (use flag = 1) or not (use flag = 0). This flag can be used to exclude outliers (e.g. due to turbidites) or radiocarbon reversals in a later analysis of the data, while maintaining the original data in the file. Isotope values without replicates were imported with a use-flag set to “1”. For replicate stable isotope measurements the use flags were set to “0” and an average of the replicates (use flag = 1) was added to the series with a comment “Mean of multiple measurements” in the same row. Raw radiocarbon ages were generally imported with a use-flag set to “0” since the data are uncalibrated. Most of the data are archived with original downcore depth of the samples. If a composite depth scale was used (e.g., for International Ocean Discovery Program (IODP) and its predecessors Deep Sea Drilling Program (DSDP) and Ocean Drilling Program (ODP) cores), a comment was added and care was taken that available radiocarbon dates were imported on the same depth scale. The data is stored as raw data, with all documented corrections removed from the data. This includes a previously subtracted reservoir age and corrections applied to the stable isotope values (e.g., to account for species offsets). Variables to store downcore radiocarbon reservoir and stable isotope corrections that may be applied to the data at a later stage are already included in the netCDF files. These variables have been imported with default values of 0.4 ka ( $\pm 0.1$ ) for all radiocarbon reservoir ages and a stable isotope correction of “0” for all oxygen and carbon isotope ratios. Both reservoir ages and stable isotope corrections can be edited within PDV.

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### 3 Data distribution

#### 3.1 Spatial and vertical coverage

Stable isotope records are available from all major ocean basins (Fig. 1), but tend to cluster along continental margins, where higher sedimentation rates, and thus higher temporal resolutions, can be found compared to mid-ocean ridges or deep abyssal basins. About 65% of the downcore records are from coring locations within 400 km of the coastline (Fig. 2). The deepest record in the atlas is from 5,105 m water depth (EN066-29PG, eastern tropical Atlantic (Curry and Lohmann, 1983), the shallowest record from 50 m water depth (GeoB9503-5, Senegal Mudbelt, Mulitza, unpublished). However, the availability of records decreases in waters shallower than about 400 m (Fig. 3), where more dynamic sedimentation regimes exist, and below 4200-3800 m due to carbonate dissolution which often prevents the production of reliable, continuous foraminiferal stable isotope records.

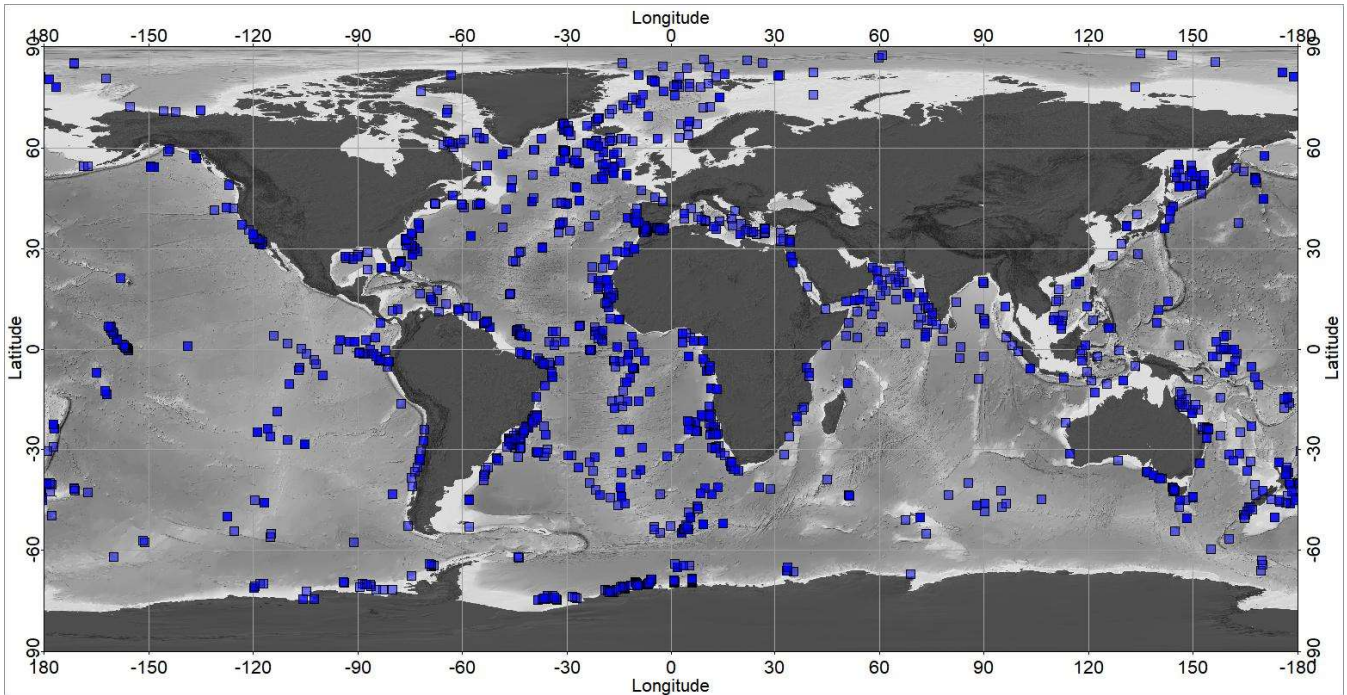
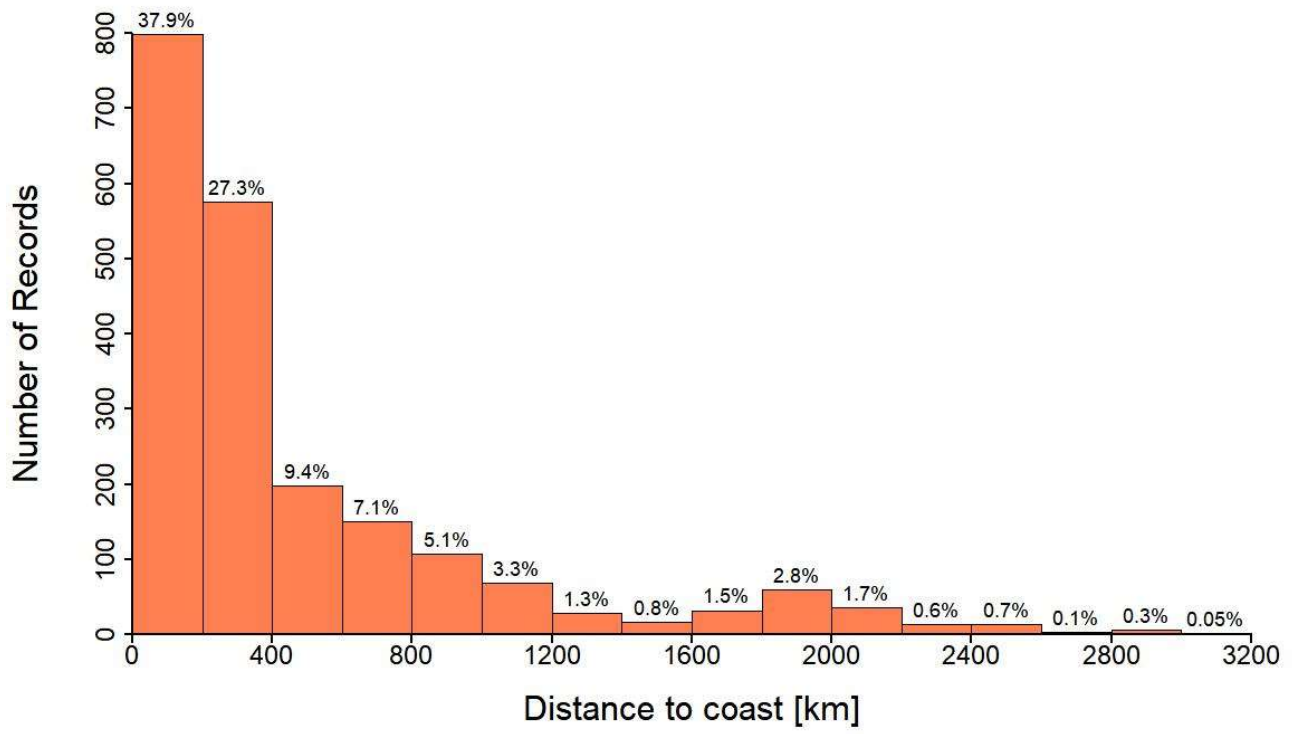


Figure 1: Spatial distribution of stable isotope records available in this atlas. The map has been generated with PaleoDataView (Langner and Mulitza, 2019).

Records are available in all oceanic 5° latitude bands with the highest number in tropical latitudes and decreasing numbers towards high latitudes (Fig. 3). This pattern is likely the result of the year-round accessibility of low latitudes compared to high latitudes where, due to sea ice cover or harsh weather conditions in the cold season, expeditions are often constrained to the warm season. The largest fraction (~47%) of the stable isotope records was measured on material from the Atlantic whereas

about 9% are from the Southern Ocean (Fig. 4). However, with 21 cores/million km<sup>2</sup>, the Mediterranean has the highest density of cores followed by the Arctic Ocean (7.8 cores/million km<sup>2</sup>) and the Atlantic (7.5 cores/million km<sup>2</sup>). The Pacific and Indian Oceans are currently only covered by 2 and 2.1 cores/million km<sup>2</sup>, respectively, which is likely a result of relatively low accumulation rates and poor carbonate preservation over large areas. In addition, the retrieval of sediment cores in the remote and deep central areas requires more ship time compared to the Atlantic and Mediterranean Sea.~~The Pacific and Indian Oceans are currently only covered by 2 and 2.1 cores/million km<sup>2</sup>.~~



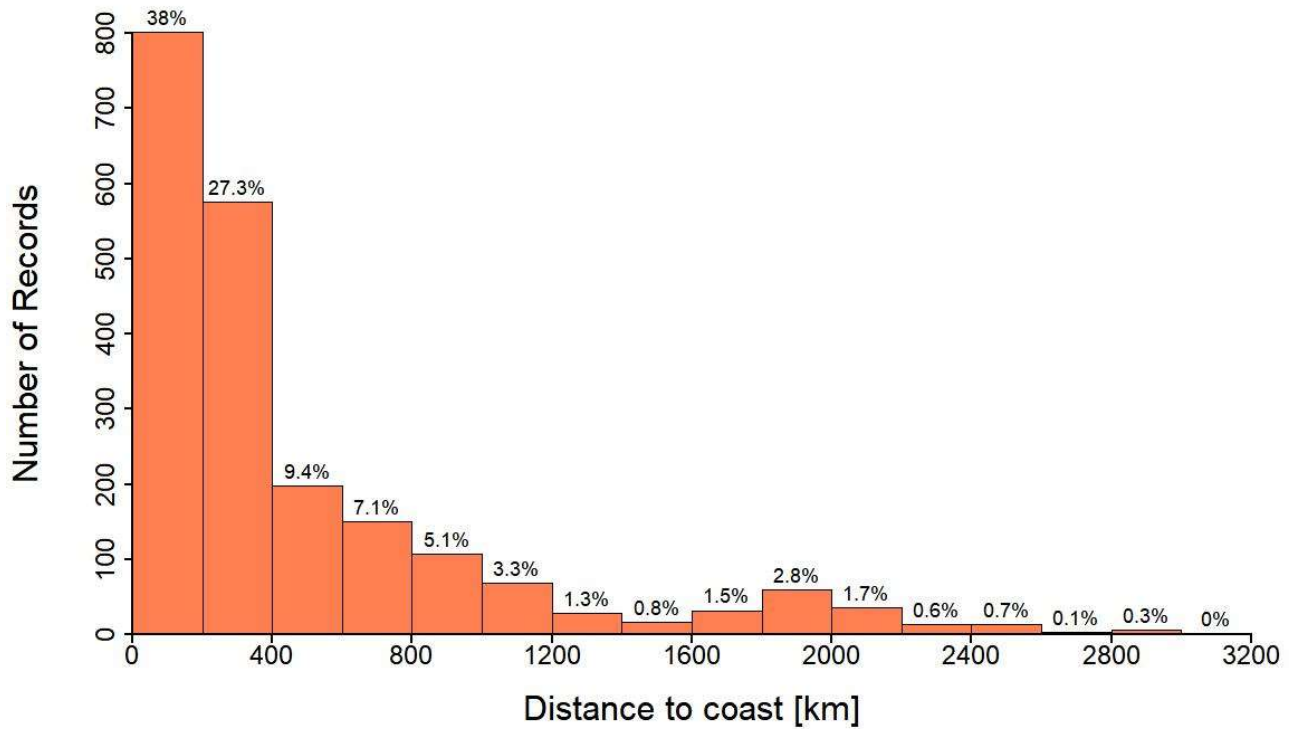


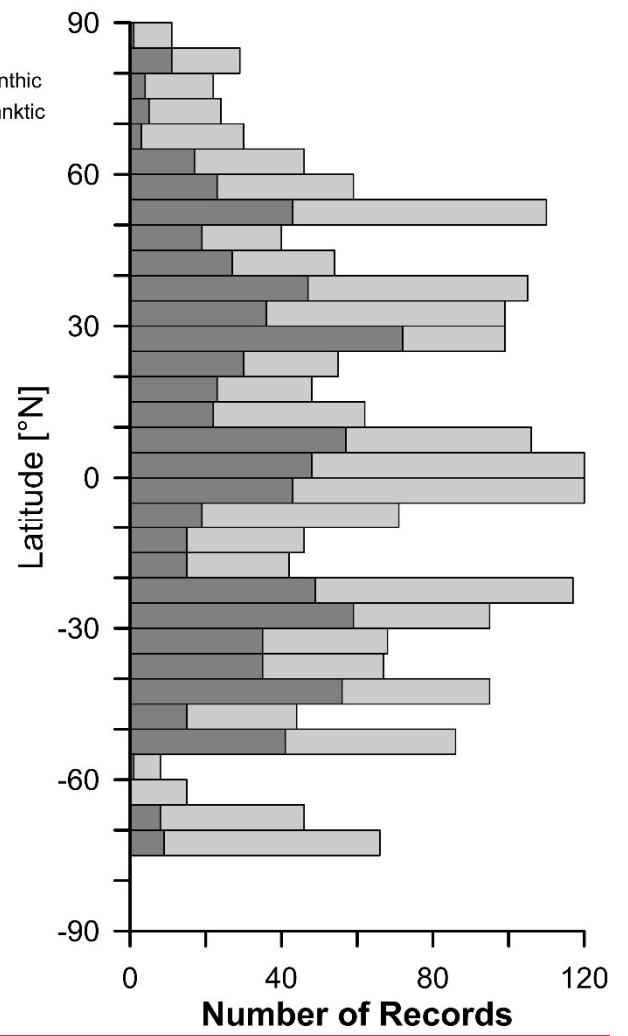
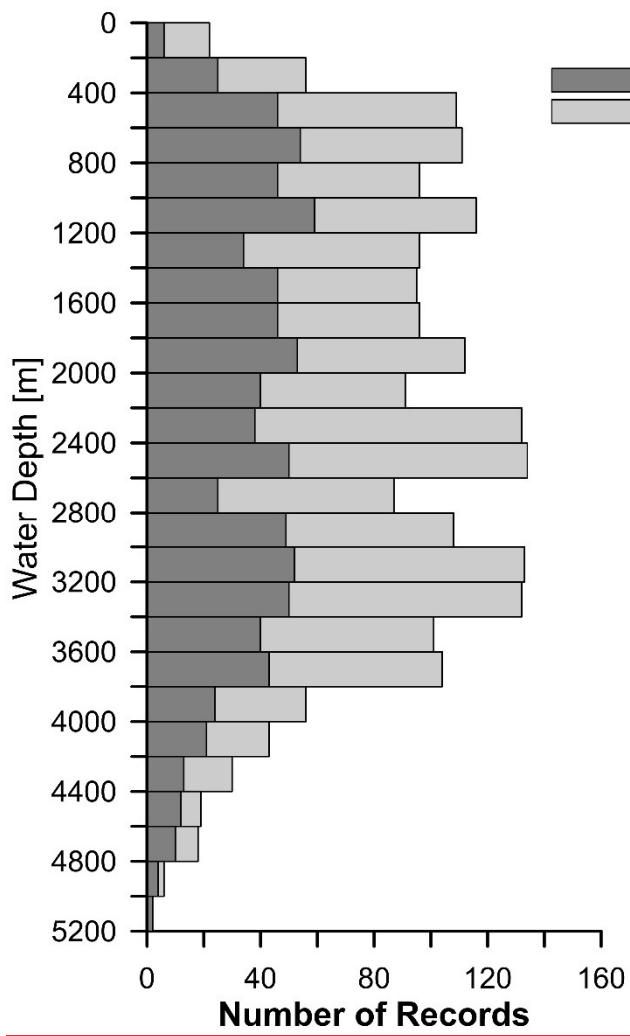
Figure 2: Number of isotope records available in this atlas versus distance to the coastline in 200 km bins. The global coastline was created with the free vector and raster map data from [www.naturalearthdata.com](http://www.naturalearthdata.com).

### 5 3.2 Species distribution

The majority (61%) of all stable isotope values available in this compilation were measured on planktic ~~foraminifera~~ Foraminifera (see individual percentages for carbon and oxygen isotopes in Fig. 5). Among the planktic species, *Globigerinoides ruber* (37%) and *Neogloboquadrina pachyderma* (28%) are the most commonly used species, followed by *Globigerina bulloides* (17%) and *Trilobatus Globigerinoides-sacculifer* (6%). These species have a relatively broad geographical coverage and are considered as mixed-layer species in their respective environment (Schiebel and Hemleben, 2017). Isotope measurements on other ~~(mostly deep-dwelling)~~ planktic species (summarized under “other planktics”) constitute about 12% of all values in the atlas (Fig. 5). 75% of the included planktic oxygen isotope values and 88% of the included benthic oxygen isotope values are reported together with the corresponding carbon isotope value. Most of the benthic isotope values (70%) were obtained from species of the epi-faunal-genus *Cibicides/Cibicidoides*. Isotope values from

the in-faunal genus *Uvigerina* constitute about 18% of all benthic isotope values in the atlas. The grouping of the original species names into species/genus names used in Fig. 5 and Fig. 6 is provided in Supplement S4.





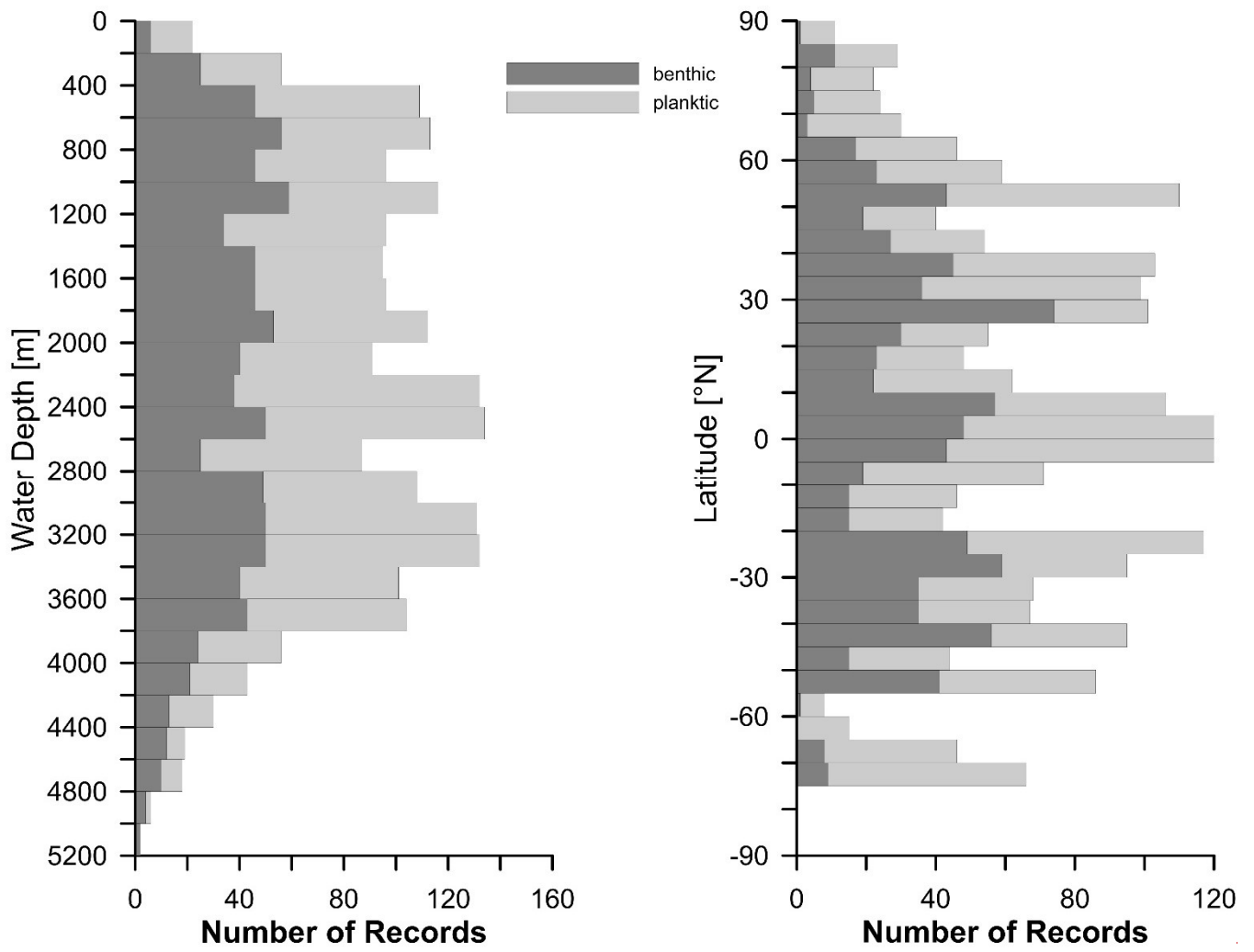
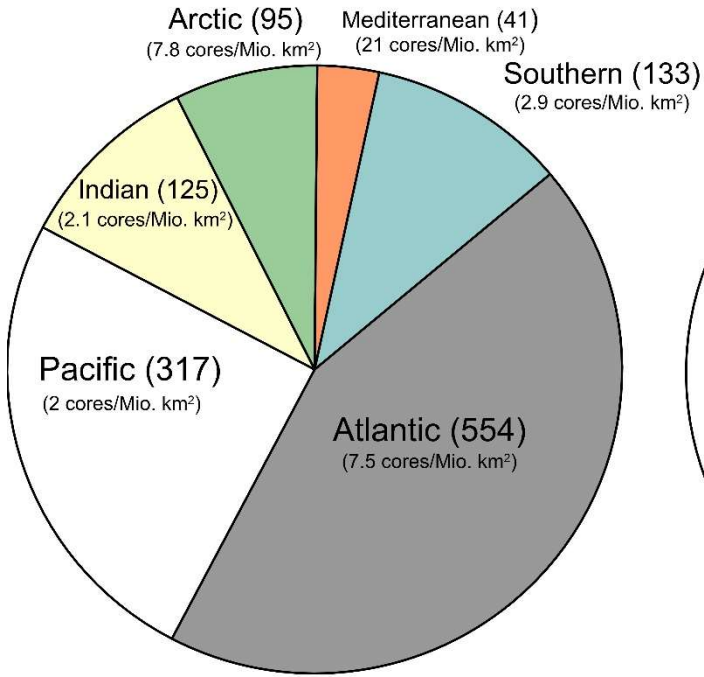
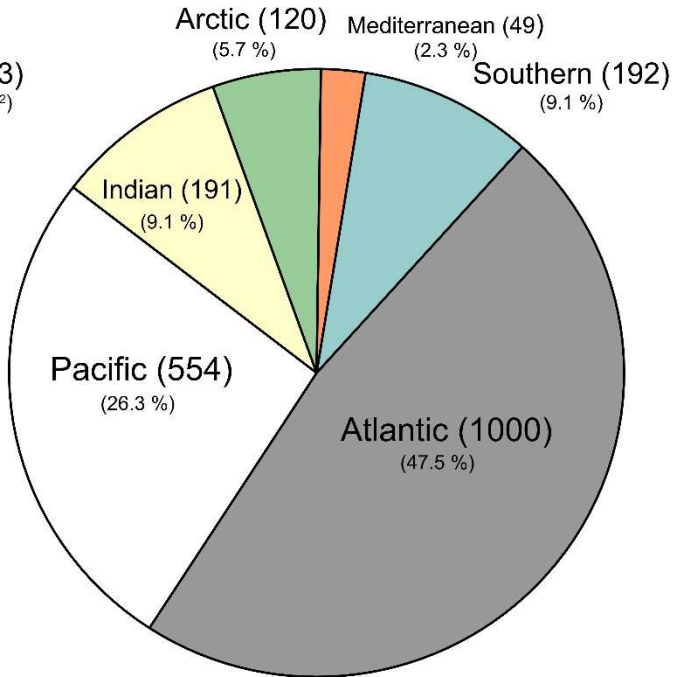


Figure 3: Distribution of stable isotope records with water depth in 200 m bins (left) and with latitude in 5° bins (right).

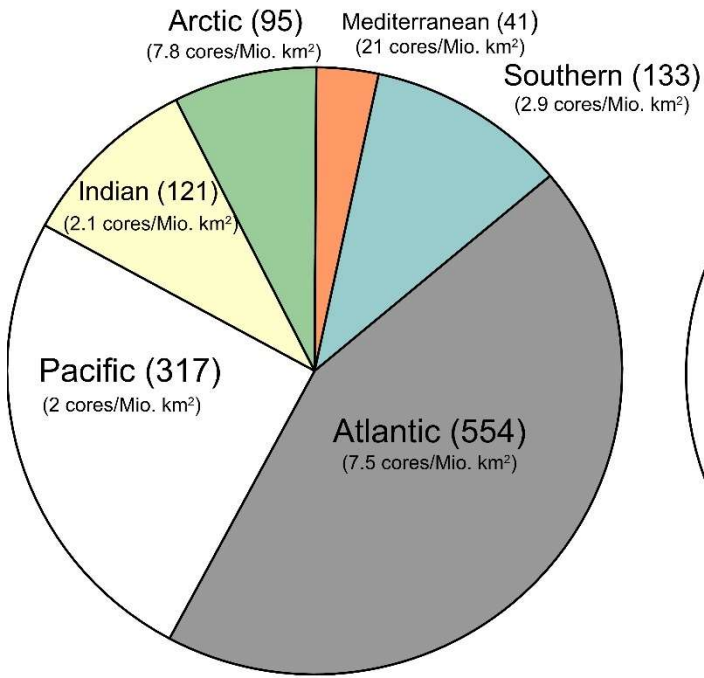
### Cores/Ocean



### Records/Ocean



### Cores/Ocean



### Records/Ocean

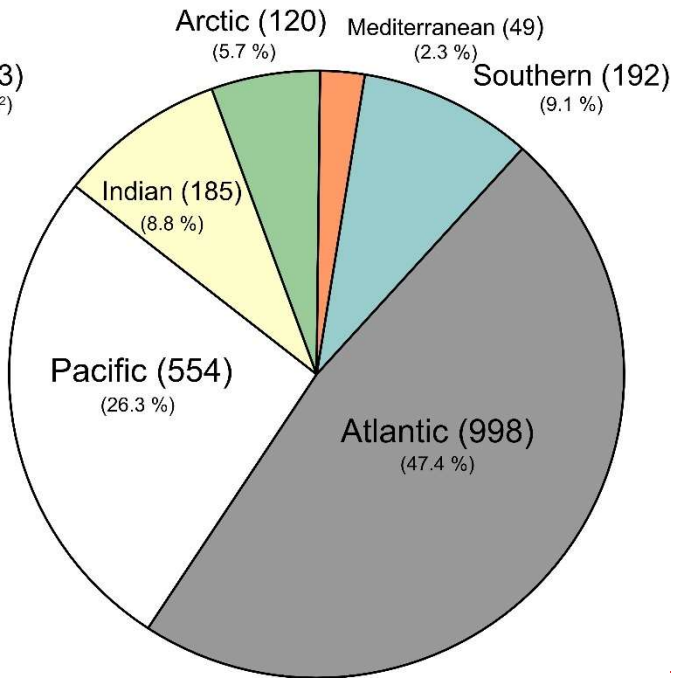
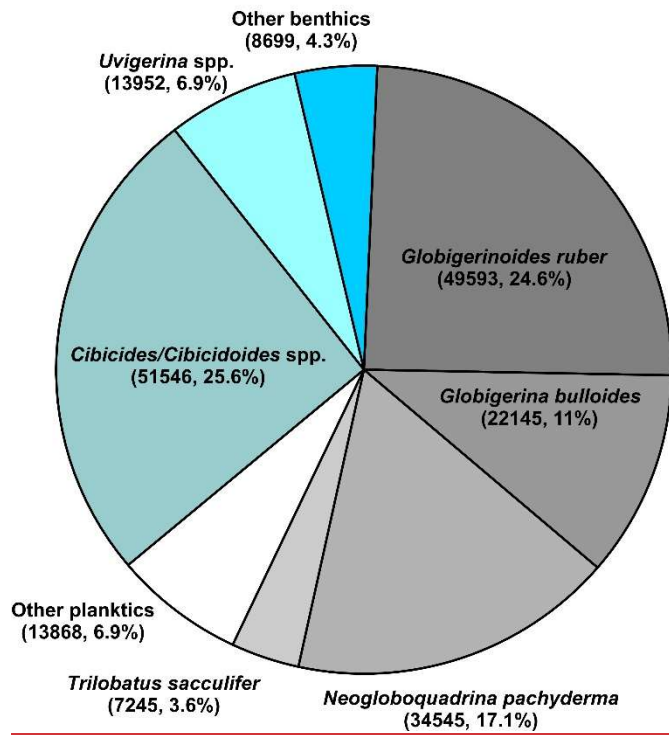


Figure 4: Number of cores (left) and records (right) for major ocean basins. A record is a downcore series of paired oxygen and carbon isotope measurements on a foraminiferal species or species group stored in a single netCDF file. Several records can exist for a single core. The counts include records/cores for which either  $\delta^{18}\text{O}$  or  $\delta^{13}\text{C}$  is missing. **Numbers in small font below ocean basin name indicate density of cores** Numbers in small letters indicate density of in cores in  $/10^6 \text{ km}^2$  (left) and percentage from the total number of records in the atlas (right) in each basin. Ocean basins follow the definitions in the World Ocean Atlas 2001 (Stephens et al., 2002). Pacific includes the Sea of Japan and the Indian Ocean includes the Bay of Bengal. ~~8 records from four and the Red Sea cores are not shown.~~

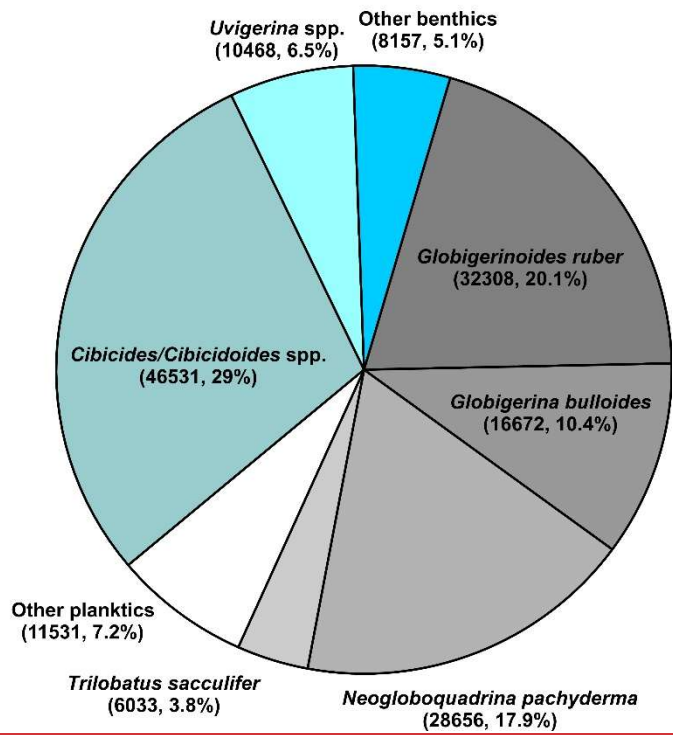
### 10 3.3 Species-specific and latitudinal distribution of oxygen and carbon isotope values

In the current version, the atlas contains a total of ~~201,593~~**201,652**  $\delta^{18}\text{O}$  values. The lowest  $\delta^{18}\text{O}$  values ( $-7.51 \text{ ‰}$ ) ~~are is~~ observed in the species *G. ruber* white (Fig. 6) from the Gulf of Mexico core LOUIS1924 (~~Aharon, 2003~~) under the influence of Mississippi freshwater discharge (~~Aharon, 2003~~). The highest planktic  $\delta^{18}\text{O}$  values ( $6.31 \text{ ‰}$ ) can be observed in the tropical species *G. ruber* from core M31\_2-78\_PC6 (Red Sea, Geiselsart and Hemleben, 1998a). With latitude, planktic  $\delta^{18}\text{O}$  values follow the typical bell-shaped curve as expected from a dominant influence of sea surface temperature (Fig. 7). Benthic  $\delta^{18}\text{O}$  values range from  $-2.85 \text{ ‰}$  from *Cibicides corpulentus* in core OC205-2-108GGC ~~from the western tropical North Atlantic~~ (Slowey and Curry, 1995) to ~~6.465.9~~  $6.465.9 \text{ ‰}$  from *Uvigerina bifurcata* ~~Creseis acieula~~ from ~~Red Sea~~South Atlantic site ~~JR244-GC528~~**M31\_2-84\_PC6**(Roberts et al., 2016). Vertically, the  $\delta^{18}\text{O}$  of ~~predominantly endobenthic and epibenthic~~ foraminiferal species increases with water depth over the upper 800 m as expected from decreasing temperatures within the main thermocline (Fig. 8). Planktic  $\delta^{18}\text{O}$  values do not show clear visual trends with water depth (not shown). Planktic and benthic  $\delta^{18}\text{O}$  values converge towards polar regions as expected from the decreasing temperature stratification with increasing latitude (Fig. 7).

$\delta^{18}\text{O}$  values/Species



$\delta^{13}\text{C}$  values/Species



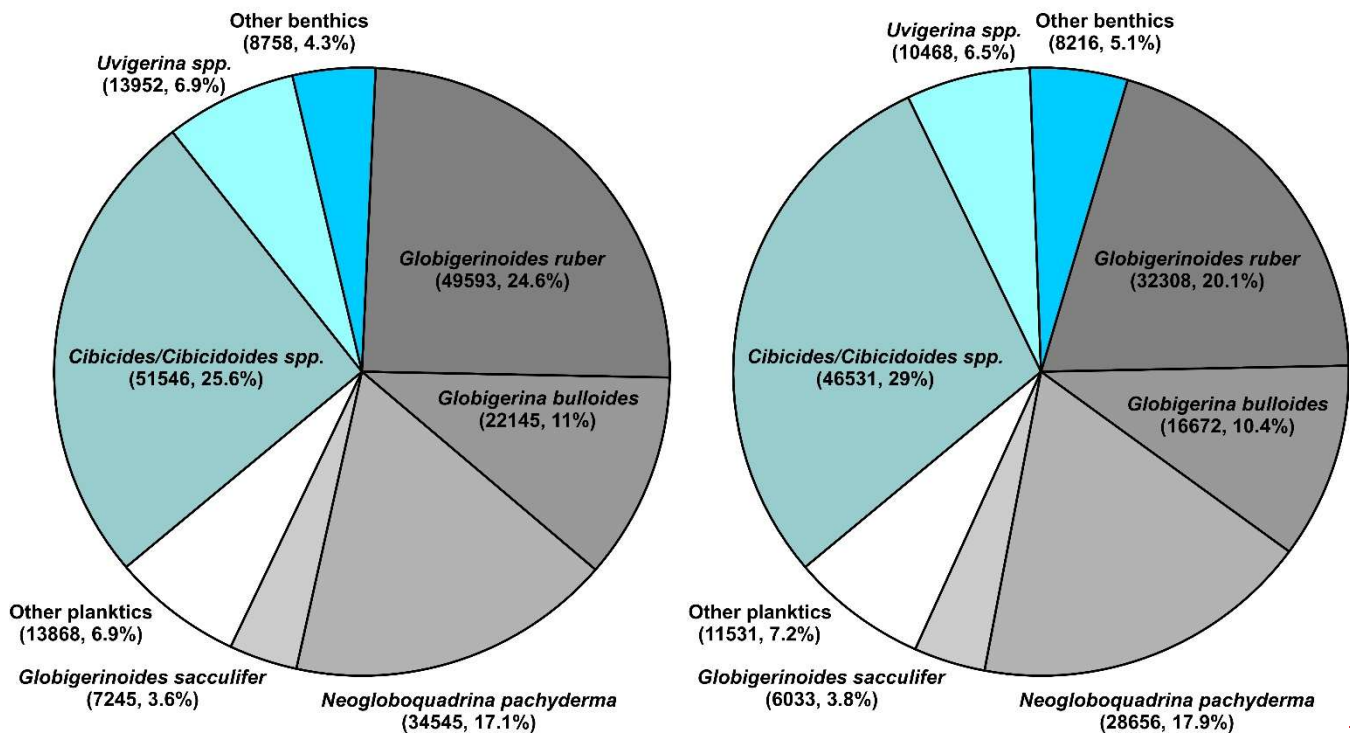
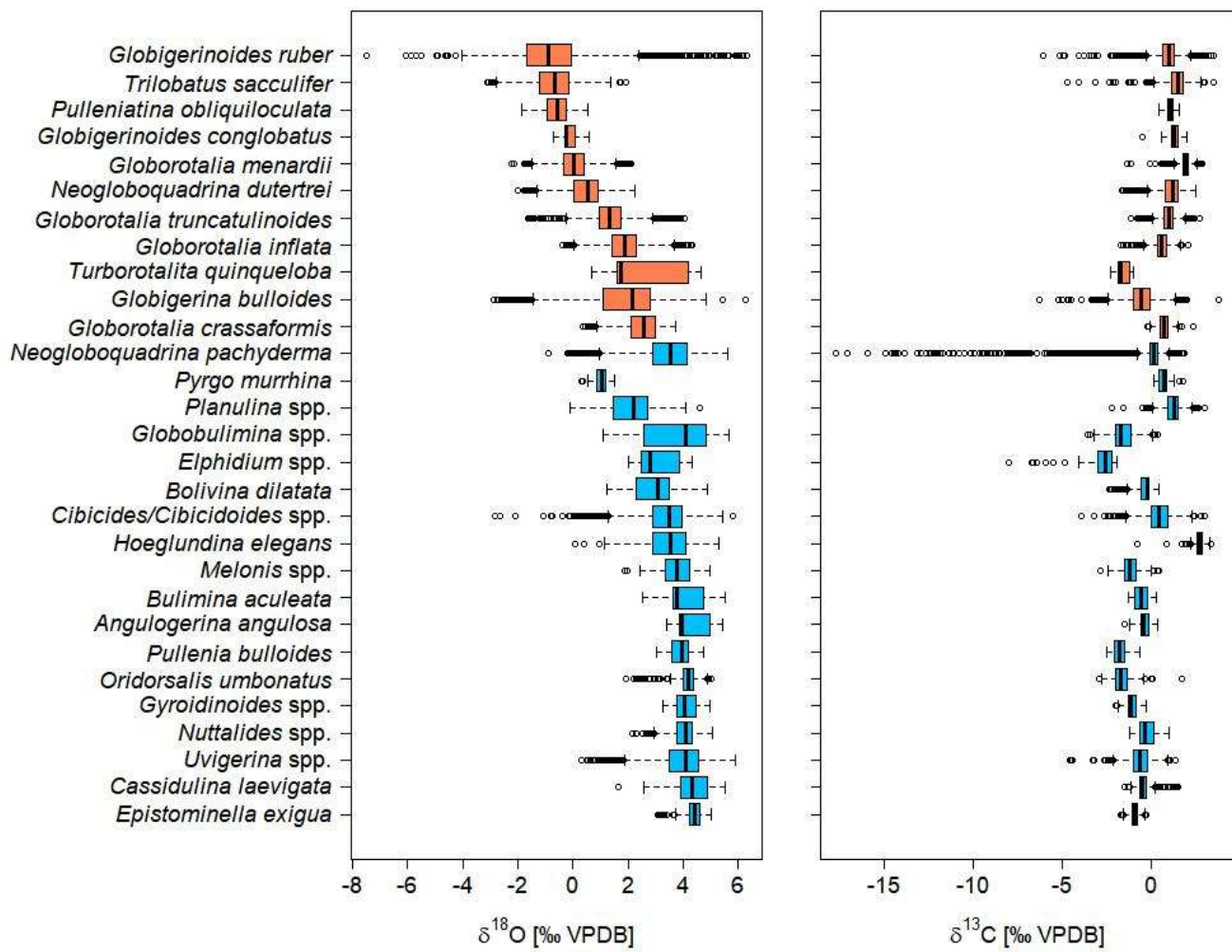
$\delta^{18}\text{O}$  values/Species $\delta^{13}\text{C}$  values/Species

Figure 5: Fraction of oxygen (left) and carbon (right) isotope values measured on benthic (blue) and planktic (grey) species/species groups. Numbers in small letters below the species/genus names indicate absolute number of values and percentage from the total number of  $\delta^{18}\text{O}$  (left) or  $\delta^{13}\text{C}$  (right) values in the atlas. See Supplement S4 for the categorisation of the individual species names from the original publications.

The data set contains 160,356~~160,415~~  $\delta^{13}\text{C}$  values. Shallow-dwelling planktic foraminifera ~~Foraminifera~~ from tropical latitudes show the highest  $\delta^{13}\text{C}$  values (Fig. 97) of up to 3.53 ‰ in shells of the species *G. ruber* from the Red Sea core M31\_2-78\_PC6 (Geiselhart and Hemleben, 1998). Planktic  $\delta^{13}\text{C}$  values get as low as -17.7 ‰ on *N. pachyderma* sinistral (Fig. 6) in core LV28-4-4 from the Sea of Okhotsk (Kaiser, 2001), which might be related to a potential contribution from authigenic carbonate minerals that form with the anaerobic oxidation of methane (Cook et al., 2011). Benthic foraminiferal  $\delta^{13}\text{C}$  gets as low as -7.99 ‰ in *Elphidium batialis* from western North Pacific core KT90-9\_21 (Oba and Murayama, 2004) and as high as 3.573.36 ‰ in the aragonitic shells of *Hoeglundina elegans*~~*Creseis aeioula*~~ from western North Atlantic~~Red-Sea~~ core OC205-2-149JPC~~M31\_2-84\_PC6~~ (Slowey and Curry, 1995). In contrast to endobenthic Benthic species, epibenthic species of the genus *Cibicides/Cibicidoides* show a clear trend toward decreasing  $\delta^{13}\text{C}$  values in the deep ocean (Fig. 8), as expected from the global distribution of  $\delta^{13}\text{C}$  in dissolved  $\Sigma\text{CO}_2$  (Kroopnick, 1985).







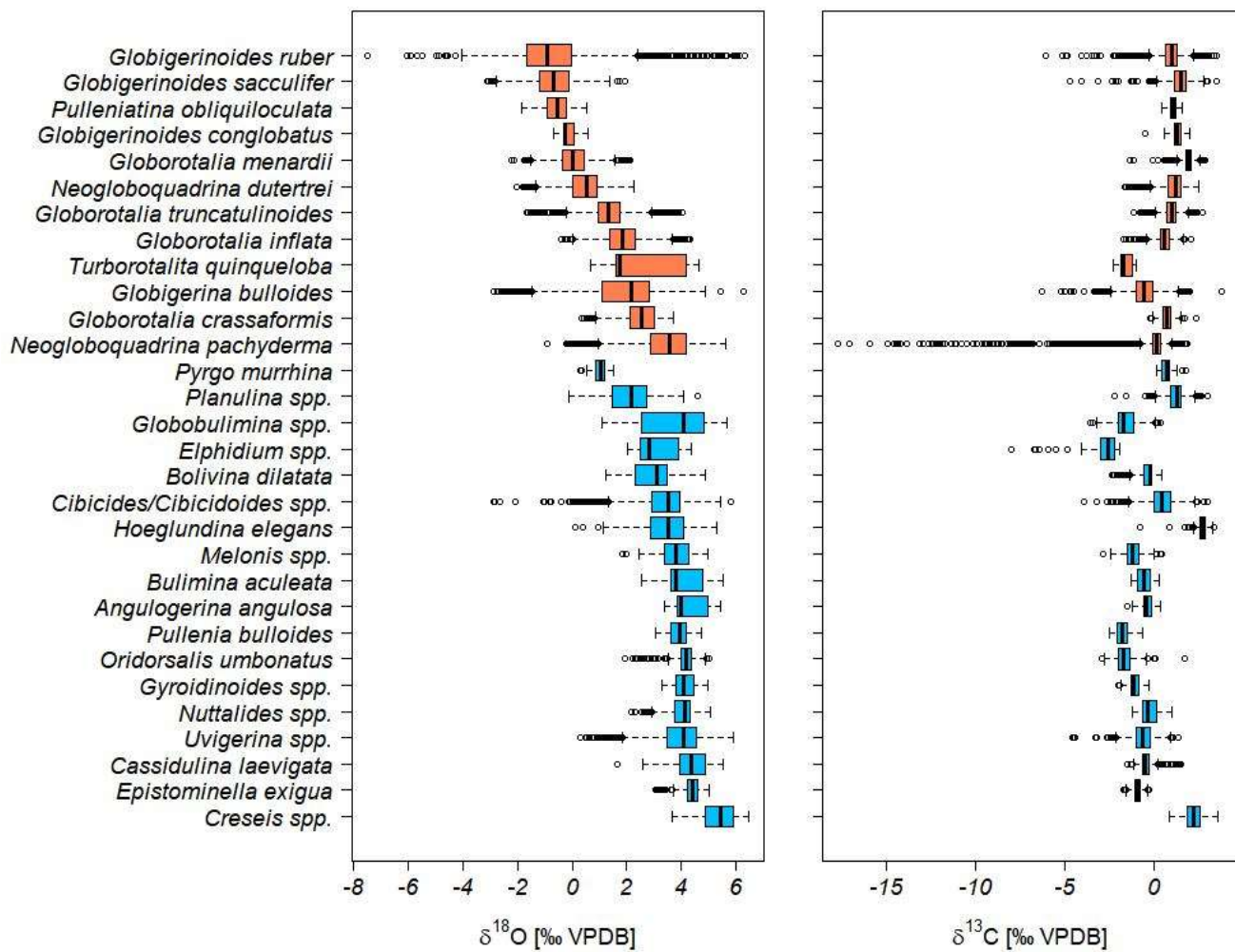
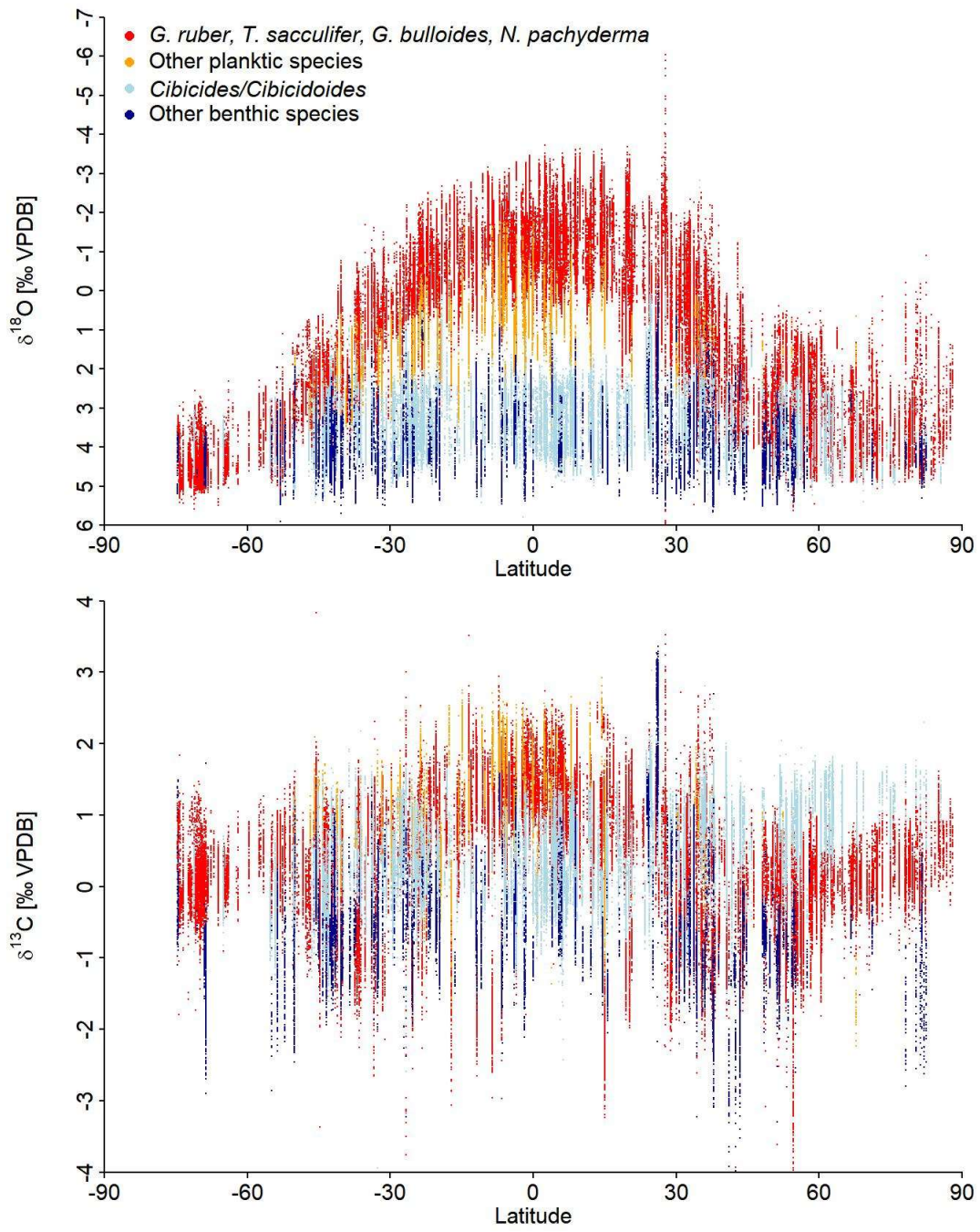


Figure 6: Box-whisker plot of oxygen (left) and carbon (right) stable isotope values of planktic (orange) and benthic (blue) **F**oraminifera at the species or genus level. The vertical line shows the median, left and right margins of the box indicate the 25<sup>th</sup> and 75<sup>th</sup> percentiles. The whiskers (the horizontal dashed lines) indicate the maximum/minimum values, or in case of outliers (open circles), highest/lowest data point that is less than 1.5 times above/below the interquartile range. The plot has been created with R's `boxplot()` function (R Core Team, 2017).



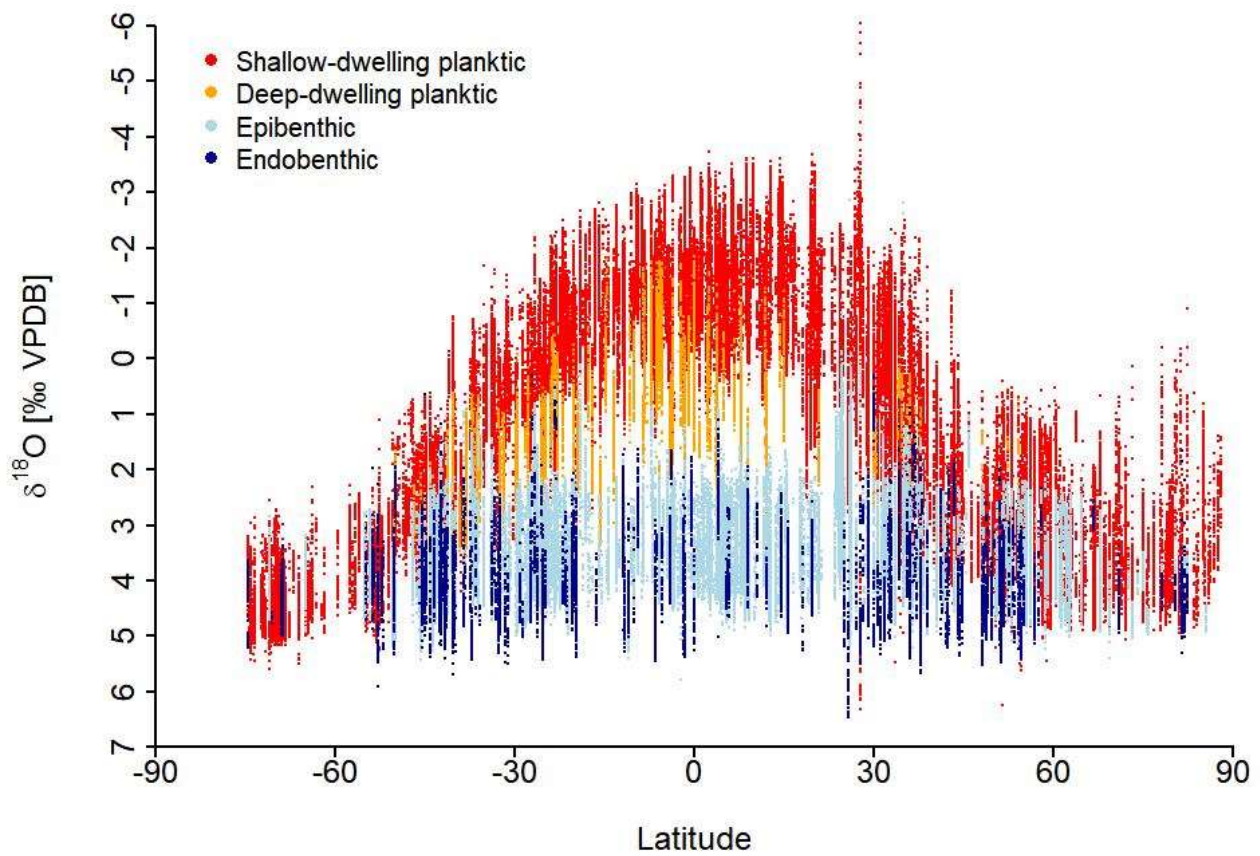
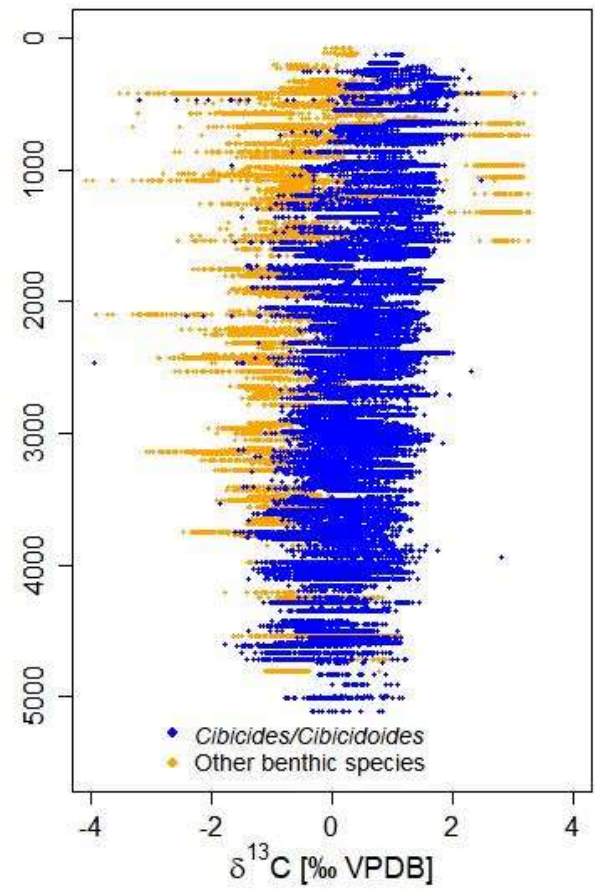
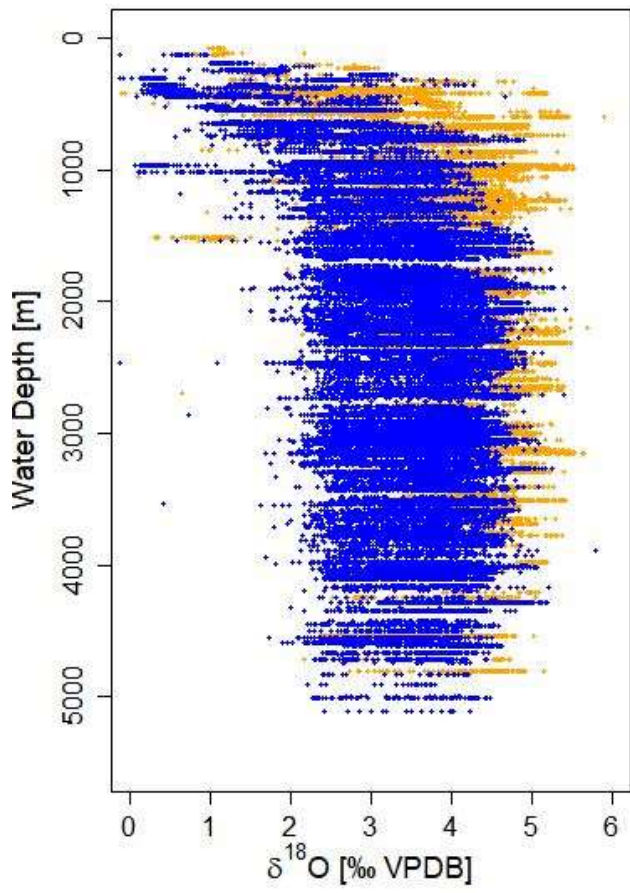


Figure 7: Distribution of  $\delta^{18}\text{O}$  values (top) and  $\delta^{13}\text{C}$  values (bottom) with latitude. Red/orange: planktic foraminifera, light/deep blue: benthic foraminifera. Extreme values outside the axis ranges are not shown.



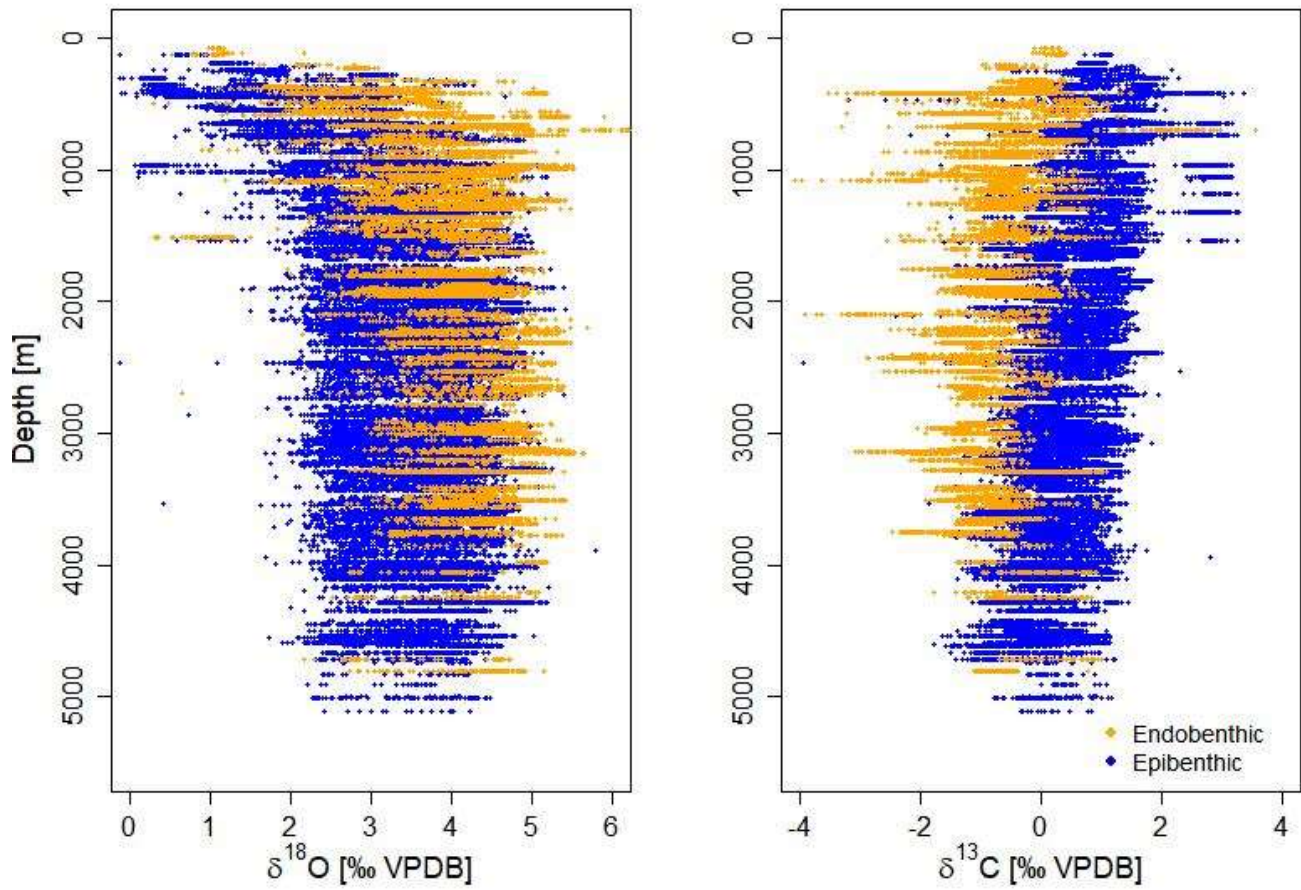


Figure 8: Distribution of benthic oxygen ( $\delta^{18}\text{O}$ , left) and carbon ( $\delta^{13}\text{C}$ , right) isotope values with water depth. Extreme values outside the axis ranges are not shown. Blue: *Cibicides/Cibicidoides*, orange: other benthic species, Orange: endobenthic species, light/deep blue: epibenthic species.

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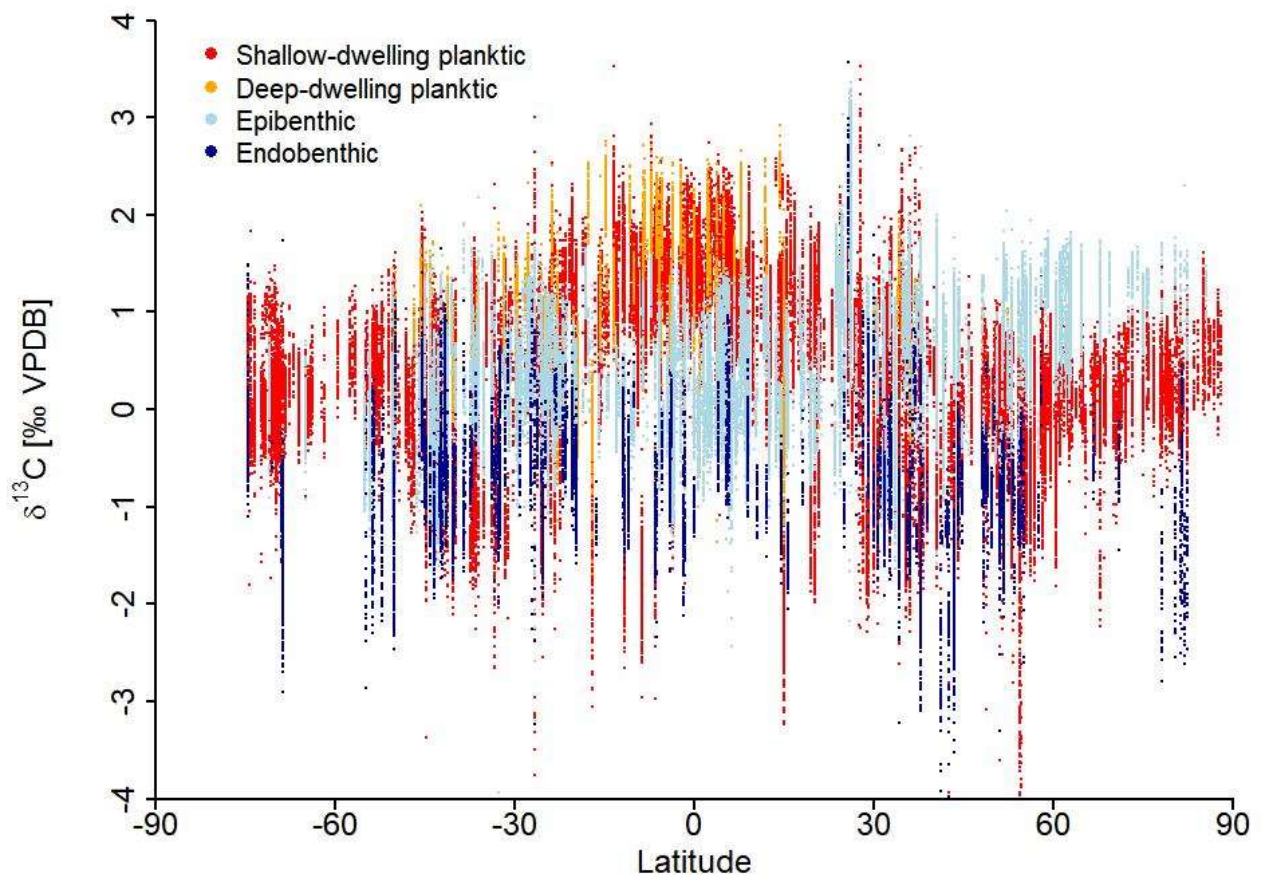


Figure 9: Distribution of  $\delta^{13}\text{C}$  values with latitude. Red/orange: planktic foraminifera, blue: benthic foraminifera. Extreme values outside the axis ranges are not shown.

### 3.4. Distribution of radiocarbon ages

The data set contains 6,153 individual radiocarbon ages with a maximum age of about 56 ka. About 47% of the cores are associated with at least one radiocarbon date. Most of the radiocarbon-dated cores are from the Atlantic (44%) followed by the Pacific (28%) and the Arctic Ocean (12%) (Fig. 109). The temporal distribution of the radiocarbon ages (Fig. 110) shows that ~~deglacial sequence~~ the last Deglaciation ~~have~~ has been preferentially dated which is likely a consequence of the stratigraphic extent and the scientific attention focussed on this time period and the limited stratigraphic extent of many coring techniques. The fraction of reversals is higher for the deglacial and glacial periods, where the higher sampling density increases the likelihood of reversals.

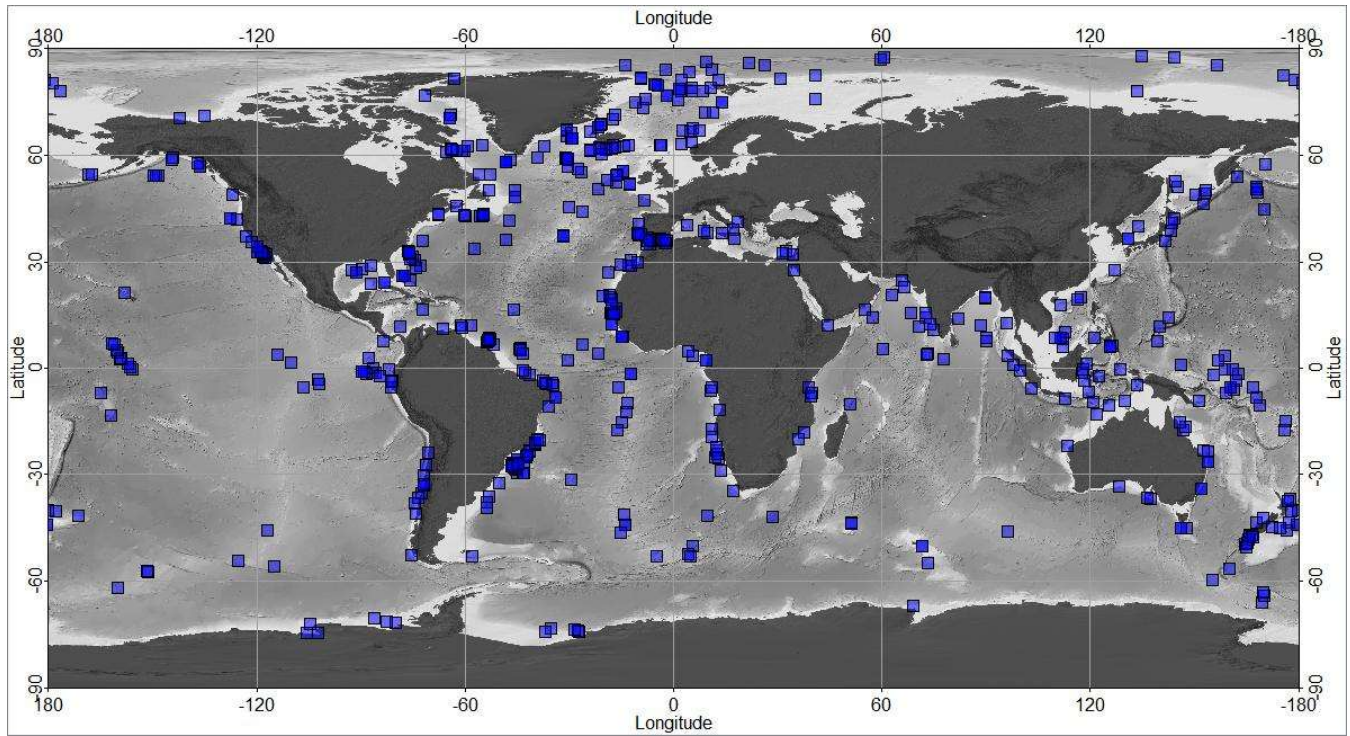


Figure 9: Spatial distribution of stable isotope records with at least one radiocarbon age. The map has been generated with PaleoDataView (Langner and Mulitza, 2019).

## 4. Possible applications

### 4.1 Marine Geology and Paleoceanography

Foraminiferal oxygen isotope ratios provide one of the most reliable tools for stratigraphy in marine sediments, particularly for time periods older than the range of the radiocarbon method, or if radiocarbon is not available or associated with large uncertainties due to unknown reservoir ages. Usually, oxygen-isotope stratigraphy is applied by using globally (Imbrie et al., 1984; Prell et al., 1986; Lisiecki and Raymo, 2005) or basin-wide (Lisiecki and Stern, 2016) isotope reference curves. The

collection presented here may provide the opportunity to find and align new records with the closest published isotope record measured on the same species, taking events into account that may only occur locally. Through its value for stratigraphy, our collection may also provide a foundation for the global mapping of seafloor sedimentation rates. The spatial quantitative mapping of sedimentation rates will allow the development of sediment budgets for the seafloor, including carbon burial.

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Oxygen and carbon isotope ratios of ~~foraminifera~~ Foraminifera are of great value for paleoclimatology, by providing information on the history of seawater temperature and isotopic composition as well as circulation, productivity and carbon sequestration. This isotope atlas will allow for new global compilations to be undertaken to understand these processes at a global scale. Although distorted by habitat- and vital effects, there is hope that some of these effects can be represented and quantified in foraminiferal ecosystem/calcification models (e.g. Wolf-Gladrow et al., 1999; Schmidt and Mulitza, 2002; Fraile et al., 2008). Since the number of climate models containing the cycling of oxygen- and carbon isotopes is constantly growing (Marchal and Curry, 2008; Kurahashi-Nakamura et al., 2017; Tierney et al., 2020; Muglia et al., 2018; Völpel et al., 2017), foraminiferal isotopes may provide the opportunity to validate climate model experiments directly. Given this prospect and the spatial coverage, foraminiferal isotope data should be rescued, assembled and organized to secure the information for future applications as ~~future work~~ we continues to improve our understanding of the ecological and geochemical processes that determine isotope ratios in foraminiferal shells. Depending on the scientific problem, paleoceanographic compilations usually have specific criteria (e.g. temporal resolution or the availability of radiocarbon ages) for the selection of the records to be included (e.g. Jonkers et al., 2020). An atlas product that includes the majority of the available records enables quick selection of suitable data without an extensive literature review.

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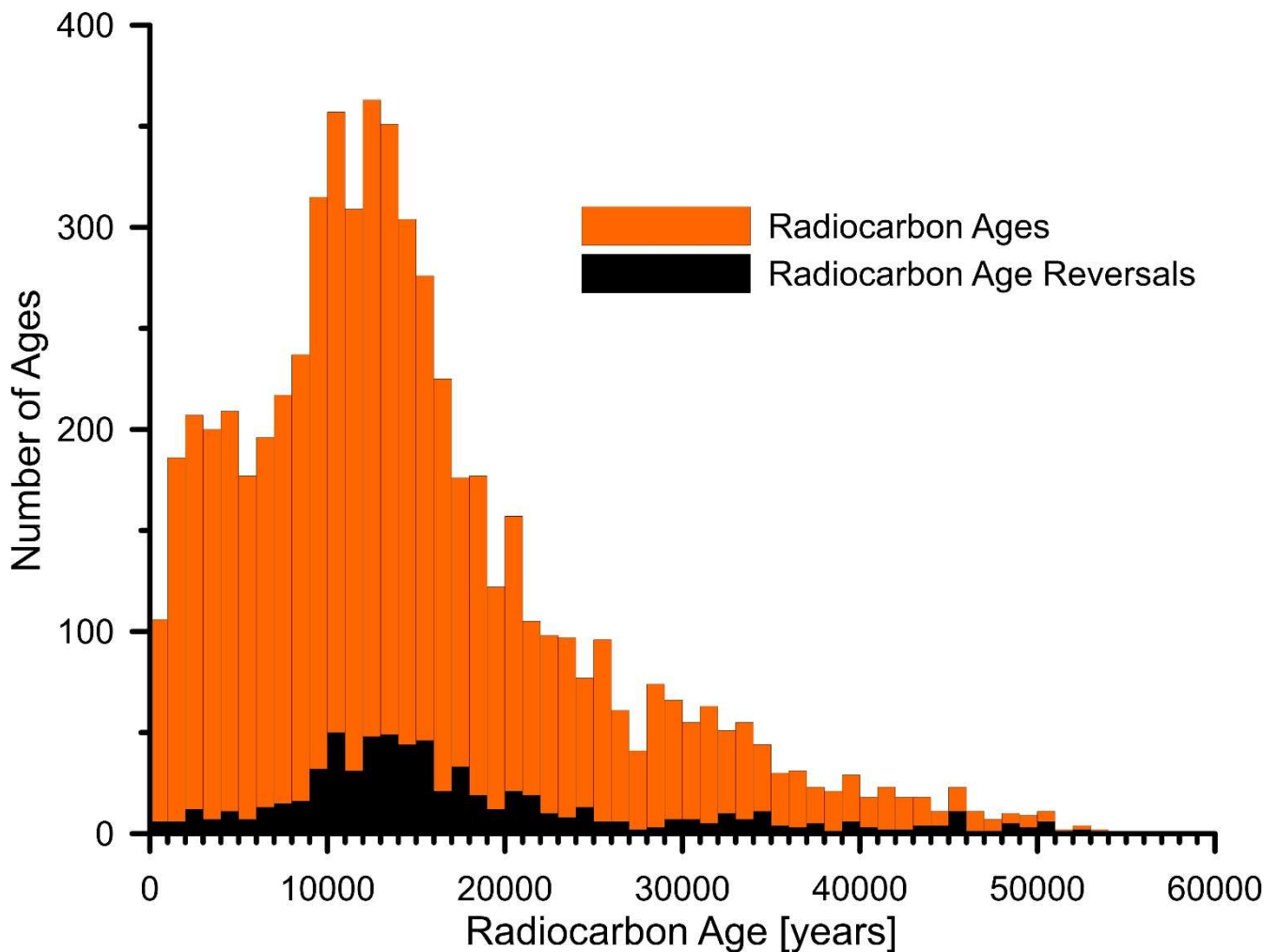


Figure 10: Distribution of radiocarbon ages in 1-kyr bins. Fraction of age reversals in black. Three negative radiocarbon ages are not included.

#### 5 4.2 Expedition planning

The planning of marine coring campaigns requires prior knowledge of existing cores. Existing core locations are often resampled to get new sediment material or to extend the stratigraphic coverage with alternative coring gear that can penetrate deeper into the sediment. For example, many IODP and ODP cores are drilled ~~on~~at sites where short cores were previously retrieved. The knowledge of existing core locations and their stratigraphy allows identification of sampling gaps. Many aspects of marine expeditions are unpredictable, and schedules and coring plans regularly have to be adapted, often on a daily basis. The atlas we are presenting here, provides fast access to stratigraphic data and may aid the identification of suitable alternative coring locations on ocean expeditions. Both the [freely available](#) PDV software and the atlas do not require web access and are therefore suitable to be used with a standard laptop computer.

### 4.3 Education

Foraminiferal oxygen isotope ratios are still the most valuable stratigraphic tool in marine sediments. The atlas covers various sedimentation regimes and therefore provides numerous examples of how factors like local hydrography, species or sedimentation rates influence the patterns of downcore isotope ratios. It therefore may be used as a resource to train students in regional isotope stratigraphy for studies in Paleoceanography, Paleoclimate and Marine Geology. Lecturers may employ the atlas together with PaleoDataView or with custom software to show examples on how isotope stages may be identified in different geological settings and on how isotope differences between species may be explained by hydrography and foraminiferal ecology. Students may also actively explore the patterns of isotope stratigraphies from different parts of the global seafloor to actively learn how global factors such as ice volume and local factors such as SST and freshwater input influence stable isotope records.

### 5 Data availability

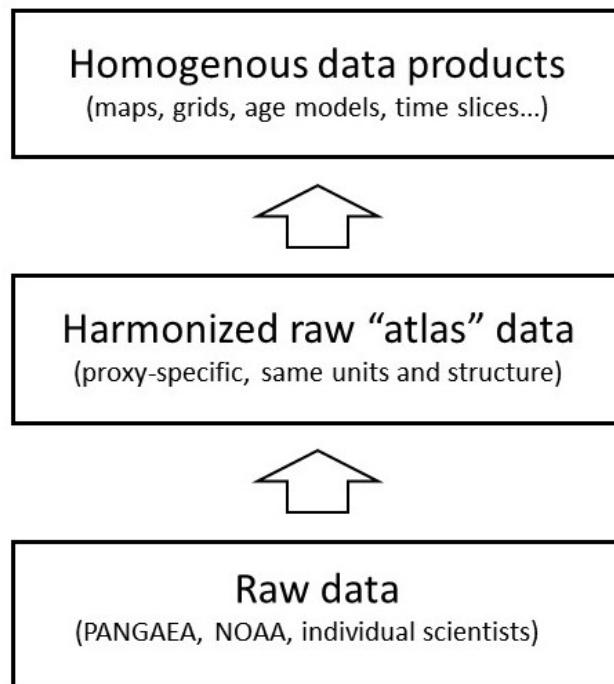
All data included in the World Atlas of late Quaternary Foraminiferal Oxygen and Carbon Isotope Ratios can be downloaded at <https://doi.pangaea.de/10.1594/PANGAEA.936747>~~https://doi.org/ (include Pangaea doi)~~. For use with the software PaleoDataView, the unzipped root directory (“WA\_Foraminiferal\_Isotopes\_2021+2022”) of the collection with all its content can be copied into the “Documents/PaleoDataView/” folder (Windows) or the /PaleoDataView/ folder under “Applications” (macOS). Select the root directory “WA\_Foraminiferal\_Isotopes\_2022” under “Data -> Change Collection -> Change Working Directory” to explore the data. For use with custom software, netCDF files containing stable isotopes data are stored under “WA\_Foraminiferal\_Isotopes\_2021+2022\Foraminiferal Isotopes\Data\” and the radiocarbon data under “WA\_Foraminiferal\_Isotopes\_2021+2022\Age\”. Installers for current versions of PDV for both Windows 10 and macOS are available from <https://www.marum.de/Stefan-Mulitza/PaleoDataView.html> (last access: March 8, 2022).

### 6 Future: Building a dynamic World Atlas of Marine Sediments

The amount of proxy data from marine sediments is growing fast and the demand of data sets that can constrain past states of the Earth system is increasing. The complexity of the data makes it challenging to maintain and reduce the data sets into spatially and chronologically coherent and meaningful data sets. We propose to initiate an atlas series that provides raw data in a consistent data format as a first step from data archived in public databases (as published) towards more sophisticated data products describing past states of the ocean and the seafloor (Fig. 1211). Eventually, these harmonized data sets can form a continuously growing and sustainable public “database layer” where proxy-specific raw data can be queried and directly loaded into software that provides the tools to generate homogenized data products that can reach out into other disciplines, i.e., climate modelling. We present a simple file-based data collection where each file contains only one proxy record rather than all available data of the core. Paleoclimatic data are often analysed and assembled in proxy-specific collections, because proxy-

specific transfer functions have to be applied in order to quantify environmental variables. Furthermore, comparisons of records from different sites are preferably done on the same proxy type to ensure comparability. A single file per proxy facilitates the composition of proxy-specific collections, avoiding the additional costs (i.e., in terms of data management and disk space) of other downcore parameters in the same file. This modularity also allows individual scientists to separate their unpublished/unvalidated data from published/validated data that are ready to be included into a proxy collection. On the other hand, it is desirable to consistently apply the same stratigraphy to all proxies from a single core. PDV will automatically apply a single age model to all proxy records with the same core label. This requires, ~~however,~~ that the depth scales and the core label of the different proxies are identical, when the data are imported.

10



**Figure 11:** Potential workflow to form sustainable data products from raw databases.

15 Stable isotopes and radiocarbon ages usually provide the stratigraphic basis for further investigations. When collections of other proxies are added to PDV, these collections can rely on the stratigraphic data provided here and any changes in the

stratigraphy will be applied to all proxy data in the collection. The efficient visualization of the data in PDV allows the identification of erroneous data and helps to improve the atlas product over time. The Excel export and import functions of PDV also ensure access to the data for individuals without strong programming skills.

- 5 [As new foraminiferal isotope measurements become frequently available, we plan to update the atlas in reasonable intervals. Also, more historical isotope data may become available and need to be rescued \(i.e. Borreggine et al., 2017\).](#) We hope this atlas will be a useful resource for the paleoceanography and marine geology community and will continue to grow through the contribution of new datasets as they are developed. [Please contact the first author if you are interested to contribute to future updates of the atlas.](#)

10

## Appendix A

Table 1: References for the included stable isotope and radiocarbon data.

<b>Core/Site</b>	<b>References</b>
12PC51	Sikes and Keigwin, 1994
3MO67	Znaidi-Rivault, 2006a
64PE-174P13	Scussolini and Peeters, 2013
75KS23	Znaidi-Rivault, 2006b
75KS5	Znaidi-Rivault, 2006c
75KS50	Znaidi-Rivault, 2006d
75KS76	Znaidi-Rivault, 2006e
75KS79	Znaidi-Rivault, 2006f
A179-15	Mix et al., 1986; CLIMAP Project Members, 2004a
A7	Sun et al., 2005
AA_GC5	Rathburn et al., 1997
AAS9_21	Govil and Naidu, 2010
AHF-11343	Mortyn et al., 1996
AHF-16830	Mortyn et al., 1996
AHF-16832	Mortyn et al., 1996
AHF-28181	Mortyn et al., 1996
AII-125JPC-76	Friddell, 2003
AII60-13APC	Curry and Lohmann, 1982
ALB226	Sarnthein et al., 1994
AMK4-316GC	Barash et al., 2002; Spielhagen et al., 1999
AOS94_B16	Poore et al., 1999
AOS94_B17	Poore et al., 1999
AOS94_B19	Poore et al., 1999
AOS94_B8	Poore et al., 1999
ASV13_1200	Duplessy et al., 2005
AT_II-107_22	Keigwin and Boyle, 1989
BA84-02PC	Kallel et al., 1997
BA84-08GC	Kallel et al., 1997
BC42-11	Showers and Margolis, 1985
BC43-15	Showers and Margolis, 1985

<b>Core/Site</b>	<b>References</b>
BC44-12	Showers and Margolis, 1985
BC5-5	Showers and Margolis, 1985
BC79-8	Showers and Margolis, 1985
BCCF10-01	Dias et al., 2018
BCCF10-01	Venancio et al., 2016
BCCF10-04	Venancio et al., 2016
BCCF10-09	Dias et al., 2018
BCCF10-15	Dias et al., 2018
BOFS14K	Bertram et al., 1995; Lowry and Machin, 2016
BOFS17K	Shimmield, 2004a
BOFS26_6K	Beveridge et al., 1995
BOFS28_3K	Beveridge et al., 1995
BOFS29_1K	Beveridge et al., 1995
BOFS30_3K	Beveridge et al., 1995
BOFS31_1K	Beveridge et al., 1995
BOFS5K	Shimmield, 2004b; Manighetti et al., 1995
BS79-33	Cacho et al., 2001; Sbaffi et al., 2001
BS88-6-10B	Horwege/Spielhagen, unpublished
BS88-6-12	Horwege/Spielhagen, unpublished
BS88-6-13	Horwege/Spielhagen, unpublished
BS88-6-14	Horwege/Spielhagen, unpublished
BS88-6-16	Horwege/Spielhagen, unpublished
BS88-6-17B	Horwege/Spielhagen, unpublished
BS88-6-18	Horwege/Spielhagen, unpublished
BS88-6-21	Horwege/Spielhagen, unpublished
BS88-6-23	Horwege/Spielhagen, unpublished
BS88-6-3	Horwege/Spielhagen, unpublished
BS88-6-4	Horwege/Spielhagen, unpublished
BS88-6-6	Horwege/Spielhagen, unpublished
BS88-6-7	Horwege/Spielhagen, unpublished
BS88-6-8	Horwege/Spielhagen, unpublished
BS-A	Ferreira et al., 2014

<b>Core/Site</b>	<b>References</b>
BS-C	Costa et al., 2018
BS-D	Ferreira et al., 2014
BT4	Curry et al., 1988
CEUTA10PC08	Ausín et al., 2015a
CF10-01B	Lessa et al., 2016; Oliveira Lessa et al., 2014
CF10-09A	Lessa et al., 2016
CH0182-36	Slowey and Curry, 1987
CH22KW31	Pastouret et al., 1978
CH69-K09	Labeyrie et al., 1999
CH71-07	Sarnthein et al., 1994
CH72-02	Curry et al., 1988
CH73-139C	Duplessy, 1982; Labeyrie and Duplessy, 1985; Bard et al., 1987
CH74-227	Labeyrie, 1996
CH75-03	Curry et al., 1988
CH75-04	Curry et al., 1988
CH84-27	Labeyrie, 1996
CHAT_16k	Yu et al., 2007
CHAT_1K	Weaver et al., 1998; McCave et al., 2008
CHAT10K	McCave et al., 2008; McCave et al., 2008; Maxson et al., 2019
CHAT3K	McCave et al., 2008
CHN115-70PC	Curry and Lohmann, 1982
CHN115-88PC	Curry and Lohmann, 1982
CHN115-89PC	Curry and Lohmann, 1982
CHN115-90PC	Curry and Lohmann, 1982
CHN115-91PC	Curry and Lohmann, 1982
CHN115-92PC	Curry and Lohmann, 1982
CHN82-20	Keigwin and Lehman, 1994
CHN82-24	Curry et al., 1988
CMU-14	Toledo et al., 2007
CS70-5	Znaidi-Rivault, 2006g
CS72-37	Kallel et al., 1997
D11957P	Lebreiro et al., 1997

<b>Core/Site</b>	<b>References</b>
DSDP590	Nelson et al., 1994, 1993
DSDP591	Nelson et al., 1994, 1993
DSDP592	Nelson et al., 1994, 1993
DSDP593	Elmore et al., 2015c
DSDP594	Nelson et al., 1986
E11-2	Mashiotta et al., 1999; Zheng et al., 2002
E27-23	Ferry et al., 2015; Anderson et al., 2009
E45-29	Howard and Prell, 1992
E49-17	Howard and Prell, 1992
E49-18	Howard and Prell, 1992
E49-21	Howard and Prell, 1992
E49-23	Howard and Prell, 1992
ELT25.011-CP	Waddell et al., 2009
ELT48.022-PC	Rickaby and Elderfield, 1999
EN066-10PG	Curry and Lohmann, 1983
EN066-16PG	Curry and Lohmann, 1983
EN066-21PG	Curry and Lohmann, 1983
EN066-26PG	Curry and Lohmann, 1983
EN066-29PG	Curry and Lohmann, 1983
EN066-32PG	Curry and Lohmann, 1983
EN066-36PG	Curry and Lohmann, 1983
EN066-38PG	Curry and Lohmann, 1983;
EN066-44PG	Curry and Lohmann, 1983
EN32-PC6	Flower et al., 2004
EN540-GGC-2	Keigwin, unpublished
ENAM9321	Rasmussen et al., 1996
ERDC-093P	Shackleton et al., 1992
ERDC-124P	Wu et al., 1990
ESP-08	Toledo et al., 2007
EW0408-26JC	Praetorius and Mix, 2014; Praetorius et al., 2015; Praetorius et al., 2016
EW0408-26TC	Praetorius et al., 2015; Praetorius et al., 2016
EW0408-66JC	Praetorius and Mix, 2014; Praetorius et al., 2016



<b>Core/Site</b>	<b>References</b>
EW0408-85JC	Praetorius et al., 2015; Davies-Walczak et al., 2014
EW0408-87JC	Praetorius et al., 2015; Davies-Walczak et al., 2014
EW9209-1JPC	Curry et al., 1999
EW9209-2JPC	Curry et al., 1999
EW9209-3JPC	Curry et al., 1999
EW9302-24GGC	Oppo et al., 2015
EW9302-25GGC	Oppo et al., 2015
EW9302-26GGC	Oppo et al., 2015
EW9504-02	Stott et al., 2000
EW9504-03	Stott et al., 2000
EW9504-04	Stott et al., 2000
EW9504-05	Stott et al., 2000
EW9504-08	Stott et al., 2000
EW9504-09	Stott et al., 2000
F2-92-P3	van Geen et al., 1996; Zheng et al., 2000
F8-90-G21	van Geen et al., 1996
Fan_17	Parker et al., 2016
FFC15	Keigwin and Lehman, 2015
FR01_97-09	Bostock et al., 2009
FR01_97-10	Bostock et al., 2009; Bostock et al., 2004
FR01_97-11	Bostock et al., 2009
FR01_97-12	Bostock et al., 2004
FR01_97-13	Bostock et al., 2009
FR01_97-14	Bostock et al., 2009
FR1_94-GC3	Deckker et al., 2019
FR4-92-PC16	Dunbar et al., 2000
FR4-92-PC36	Dunbar et al., 2000
FR4-92-PC42	Dunbar et al., 2000
FR4-92-PC6	Dunbar et al., 2000
FR5-90-PC27a	Bostock et al., 2006
GC34	Moy et al., 2006
GeoB10038-4	Mohtadi et al., 2010a; Mohtadi et al., 2010b; Mohtadi, unpublished

<b>Core/Site</b>	<b>References</b>
GeoB10053-7	Mohtadi et al., 2011; Mohtadi, unpublished
GeoB10069-3	Gibbons et al., 2014; Mohtadi, unpublished
GeoB1007-4	Mulitza and Rühlemann, 2000; Mulitza, unpublished
GeoB1008-3	Schneider, 1991; Govin, unpublished
GeoB1016-3	Schneider et al., 1995; Govin, unpublished
GeoB1023-5	Schneider et al., 1995; Kim and Schneider, 2003
GeoB1028-5	Wefer et al., 1996; Bickert and Mackensen, 2004
GeoB1031-4	Wefer et al., 1996; Bickert and Mackensen, 2004
GeoB1032-2	Bickert and Mackensen, 2004
GeoB1032-3	Wefer et al., 1996; Bickert and Mackensen, 2004
GeoB1034-1	Bickert and Mackensen, 2004
GeoB1034-3	Bickert and Mackensen, 2004
GeoB1035-3	Bickert and Mackensen, 2004
GeoB1035-4	Bickert and Mackensen, 2004
GeoB1041-1	Bickert and Mackensen, 2004
GeoB1041-3	Wolff, 1998; Bickert and Mackensen, 2004
GeoB1101-4	Bickert and Mackensen, 2004
GeoB1101-5	Bickert and Mackensen, 2004
GeoB1105-3	Kemle-von Mücke, 1994; Bickert and Mackensen, 2004
GeoB1105-4	Meinecke, 1992; Kemle-von Mücke, 1994; Bickert and Mackensen, 2004
GeoB1112-3	Bickert and Mackensen, 2004
GeoB1112-4	Kemle-von Mücke, 1994; Bickert and Mackensen, 2004
GeoB1113-4	Sarnthein et al., 1994
GeoB1113-7	Sarnthein et al., 1994
GeoB1115-3	Bickert and Mackensen, 2004
GeoB1115-4	Bickert and Mackensen, 2004
GeoB1117-2	Bickert and Mackensen, 2004
GeoB1117-3	Bickert and Mackensen, 2004
GeoB1118-2	Bickert and Mackensen, 2004
GeoB1118-3	Bickert and Mackensen, 2004
GeoB1211-1	Bickert and Mackensen, 2004
GeoB1211-3	Bickert and Mackensen, 2004

<b>Core/Site</b>	<b>References</b>
GeoB1214-1	Bickert and Mackensen, 2004
GeoB1214-2	Bickert and Mackensen, 2004
GeoB1220-1	Wefer et al., 1996; Bickert and Mackensen, 2004
GeoB12605-3	Kuhnert et al., 2014; Kuhnert, unpublished
GeoB12615-4	Romahn et al., 2014
GeoB12624-1	Liu et al., 2016; Bouimetarhan et al., 2015
GeoB1306-1	Bickert and Mackensen, 2004
GeoB1306-2	Bickert and Mackensen, 2004
GeoB1309-2	Hale and Pflaumann, 1999a
GeoB1312-2	Hale and Pflaumann, 1999a; Bickert and Mackensen, 2004
GeoB13601-4	Just et al., 2012, N. Syring 2011
GeoB13731-1	Fink et al., 2013; Wang et al., 2019
GeoB13801-2	Bender et al., 2013
GeoB13825-2	Bickert, unpublished
GeoB13862-1	Voigt et al., 2015
GeoB1408-3	Dürkoop et al., 1997a; Mulitza, 2009a
GeoB1413-4	Wefer et al., 1996
GeoB1417-1	Meinecke, 1992; Bickert and Mackensen, 2004
GeoB1419-2	Bickert and Mackensen, 2004
GeoB15005-1	Martínez-Méndez et al., 2013
GeoB1501-4	Dürkoop et al., 1997b; Bickert and Mackensen, 2004
GeoB1503-1	Dürkoop et al., 1997c; Bickert and Mackensen, 2004; Mulitza, unpublished
GeoB1505-2	Bickert and Mackensen, 2004
GeoB1506-2	Wolff, 1998
GeoB1508-4	Dürkoop et al., 1997d; Bickert and Mackensen, 2004
GeoB1515-1	Rühlemann et al., 1996; Vidal et al., 1999
GeoB1520-1	Bickert and Mackensen, 2004
GeoB1520-2	Bickert and Mackensen, 2004
GeoB1523-1	Mulitza, 1994; Bickert and Mackensen, 2004; Mulitza, 2009b
GeoB1523-1	Rühlemann et al., 2001
GeoB1523-2	Bickert and Mackensen, 2004

<b>Core/Site</b>	<b>References</b>
GeoB16202-2	Freytmüller, 2013; Vahlenkamp, 2013; Huppertz, 2014; Mulitza et al., 2017; Voigt et al., 2017; Venancio et al., 2018; Mulitza, unpublished; Mulitza and Mackensen, unpublished
GeoB16206-1	Zhang et al., 2015; Voigt et al., 2017
GeoB16224-1	Krummrei, 2015; Zhang et al., 2015; Voigt et al., 2017; Crivellari et al., 2018; Mulitza, unpublished
GeoB16320-2	Matos et al., 2017
GeoB1701-4	Dürkoop et al., 1997e; Mulitza, unpublished
GeoB1704-4	Mollenhauer, 2002; Mollenhauer, unpublished
GeoB1706-2	Little et al., 1997
GeoB1710-2	Bickert and Mackensen, 2004
GeoB1710-3	Schmiedl and Mackensen, 1997
GeoB1711-4	Bickert and Mackensen, 2004; Vidal et al., 1999; Little et al., 1997; Balmer et al. 2016
GeoB1711-5	Bickert and Mackensen, 2004
GeoB1712-4	Mollenhauer, unpublished
GeoB1720-2	Dickson et al., 2009
GeoB1721-4	Bickert and Mackensen, 2004
GeoB1721-7	Bickert and Mackensen, 2004
GeoB1722-1	Bickert and Mackensen, 2004
GeoB1722-3	Bickert and Mackensen, 2004
GeoB1903-3	Dürkoop et al., 1997f; Bickert and Mackensen, 2004; Niebler and Mulitza, 2009
GeoB1905-3	Bickert and Mackensen, 2004
GeoB2004-2	Bickert and Mackensen, 2004; Mulitza, 2009c
GeoB2016-1	Niebler, 2004g; Bickert and Mackensen, 2004
GeoB2019-1	Bickert and Mackensen, 2004; Niebler, 2004h; Mulitza, unpublished
GeoB2021-5	Niebler, 2004i; Mulitza, unpublished
GeoB2104-3	Steinborn, 2003; Hickey, 2010; Mulitza, unpublished
GeoB2105-1	Steinborn, 2003
GeoB2106-3	Steinborn, 2003
GeoB2107-3	Dürkoop, 1998; Portilho-Ramos et al., 2018; Heil, 2006; Rühlemann, unpublished
GeoB2109-1	Hale and Pflaumann, 1999b; Dürkoop et al., 2004a; Mulitza, 2009d; Huang, 2013

<b>Core/Site</b>	<b>References</b>
GeoB2110-4	Gingele et al., 1999
GeoB2116-4	Niebler, 2004j; Mulitza, 2004
GeoB2117-1	Dürkoop et al., 1997g
GeoB2125-1	Dürkoop et al., 1997h
GeoB2126-3	Govin, unpublished
GeoB2202-4	Dürkoop et al., 1997i
GeoB2204-1	Dürkoop, 1998; Bickert and Mackensen, 2004; Mulitza, unpublished
GeoB2204-2	Dürkoop, 1998; Bickert and Mackensen, 2004; Mulitza, unpublished
GeoB2215-10	Wolff, 1998; Bickert and Mackensen, 2004
GeoB2819-1	Dürkoop et al., 1997j; Hale and Pflaumann, 1999a; Bickert and Mackensen, 2004
GeoB3004-1	Schmiedl and Mackensen, 2006
GeoB3005-1	Müller and Budziak, 2004
GeoB3104-1	Arz et al., 1998; Arz et al., 1999b
GeoB3117-1	Arz et al., 1999a
GeoB3129-1	Arz et al., 1999a
GeoB3176-1	Arz et al., 1999a
GeoB3202-1	Arz et al., 1999b; Behling et al., 2002
GeoB3229-2	Arz et al., 1999b
GeoB3302-1	Lamy, 1998; Mohtadi et al., 2008
GeoB3304-5	Bernhardt et al., 2016; Bernhardt et al., 2015
GeoB3313-1	Lamy et al., 2002
GeoB3327-5	Ho et al., 2012
GeoB3359-3	Mohtadi et al., 2008
GeoB3369-1	Bernhardt et al., 2016; Bernhardt et al., 2015
GeoB3375-1	Lamy et al., 1998; Lamy et al., 2000
GeoB3603-2	Bickert and Mackensen, 2004
GeoB3606-1	Romero et al., 2003
GeoB3722-2	Mollenhauer, 2002; Niebler et al., 2003; Mollenhauer, unpublished; Niebler, unpublished
GeoB3801-6	Bickert and Mackensen, 2004; Mulitza, 2009e
GeoB3808-6	Hale and Pflaumann, 1999c; Bickert and Mackensen, 2004; Dürkoop et al., 2004b
GeoB3813-3	Bickert and Mackensen, 2004; Mulitza, 2009f

<b>Core/Site</b>	<b>References</b>
GeoB3914-2	Govin, unpublished
GeoB3935-2	Schlünz et al., 2000
GeoB3938-1	Schlünz et al., 2000; Govin et al., 2014a
GeoB4216-1	Freudenthal et al., 2002
GeoB4223-2	Freudenthal et al., 2002; Henderiks et al., 2002
GeoB4240-2	Freudenthal et al., 2002; Henderiks et al., 2002
GeoB4241-11	Freudenthal, 2000; Henderiks et al., 2002
GeoB4403-2	Bickert and Mackensen, 2004
GeoB4411-2	Hörner, 2012; Govin et al., 2014a
GeoB4420-2	Mulitza, unpublished
GeoB4901-8	Adegbie, 2001
GeoB4905-4	Adegbie et al., 2003; Weldeab et al., 2005; Zimmermann, 2013
GeoB5115-2	Niebler, 2004k; Bickert and Mackensen, 2004
GeoB5121-2	Niebler, 2004l; Bickert and Mackensen, 2004
GeoB5844-2	Arz et al., 2003
GeoB5901-2	Schirmacher et al., 2020; Rühlemann, unpublished
GeoB6201-5	Portilho-Ramos et al., 2018
GeoB6211-2	Steinborn, 2003; Chiessi et al., 2008; Chiessi et al., 2009; Voigt et al., 2015; Chiessi, unpublished
GeoB6212-1	Chiessi and Mulitza, unpublished
GeoB6213-2	Mulitza and Chiessi, unpublished
GeoB6308-3	Voigt et al., 2015
GeoB6340-2	Mulitza, unpublished
GeoB6403-3	Donner, unpublished
GeoB6405-6	Donner, unpublished
GeoB6408-4	Donner, unpublished
GeoB6412-2	Barbara Donner, unpublished
GeoB6421-2	Barbara Donner, unpublished
GeoB6425-2	Donner, unpublished
GeoB6518-1	Schefuss et al., 2005
GeoB6719-1	Rüggeberg et al., 2005
GeoB6910-2	Steinborn, 2003

<b>Core/Site</b>	<b>References</b>
GeoB6914-2	Steinborn, 2003
GeoB7010-2	Kuhr, 2011 Govin et al., 2014b; Govin et al., 2014a
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GIK16776-1	Hüls, 1991
GIK16856-2	Sarnthein et al., 1994; Schulz, 1995
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GIK17049-6	Jung, 1996
GIK17050-1	Jung and Sarnthein, 2003a
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GIK17304-1	Winn, 2013e
GIK17304-2	Winn, 2013f
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GIK17748-2	Mohtadi and Hebbeln, 2004; Mohtadi et al., 2008
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GIK23519-4	Millo et al., 2006
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GL-74	Portilho-Ramos et al., 2014
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GS07-150_11_1MC-C	Santos et al., 2013
GS07-150_17_2MC-A	Santos et al., 2013
GS07-150_MC-B	Santos et al., 2014
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HLY02-02-51	Cook et al., 2011; Caissie et al., 2010
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KF-12	Costa et al., 2016a
KF14	Leonhardt et al., 2015
KF16	Repschläger et al., 2015
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KN07304-0003PG	Curry et al., 1988
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KN166-14-JPC-13	Hodell et al., 2010
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KNR110-50	Curry et al., 1988
KNR110-55	Sarnthein et al., 1988
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KNR110-75	Curry et al., 1988
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KNR140-02JPC	Keigwin, 2004
KNR140-02PG	Keigwin, 2004
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KNR159-5-120GGC	Hoffman and Lund, 2012
KNR159-5-125GGC	Lund et al., 2015; Hoffman and Lund, 2012
KNR159-5-14GGC	Lund et al., 2015
KNR159-5-17JPC	Lund et al., 2015; Tessin and Lund, 2013
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KNR159-5-22GGC	Lund et al., 2015; Hoffman and Lund, 2012
KNR159-5-30GGC	Lund et al., 2015; Tessin and Lund, 2013
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KNR159-5-36GGC	Lund et al., 2015; Carlson et al., 2008; Sortor and Lund, 2011; Came et al., 2003
KNR159-5-42JPC	Lund et al., 2015; Hoffman and Lund, 2012
KNR159-5-54GGC	Hoffman and Lund, 2012
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KNR159-5-78GGC	Lund et al., 2015; Tessin and Lund, 2013
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KNR166-2-105JPC	Lynch-Stieglitz et al., 2009
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KNR166-2-135JPC	Lynch-Stieglitz et al., 2009
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M35027-1	Stüber, 1999
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MD76-127	Sirocko, 1989
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MD76-135	Sirocko, 1989
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MD81-BC15	Thunell, 2006b
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MG237	Giresse et al., 1982; Sarnthein et al., 1994
ML1208-06BB	Lynch-Stieglitz et al., 2015
ML1208-10GC	Lynch-Stieglitz et al., 2015
ML1208-11GC	Lynch-Stieglitz et al., 2015
ML1208-12GC	Lynch-Stieglitz et al., 2015
ML1208-13BB	Monteagudo et al., 2021; Lynch-Stieglitz et al., 2015; Costa and McManus, 2017
ML1208-15GC	Lynch-Stieglitz et al., 2015
ML1208-17PC	Lynch-Stieglitz et al., 2015
ML1208-17TC	Lynch-Stieglitz et al., 2015
ML1208-18GC	Lynch-Stieglitz et al., 2015; Monteagudo et al., 2021; Lynch-Stieglitz, unpublished
ML1208-19GC	Lynch-Stieglitz et al., 2015
ML1208-20BB	Monteagudo et al., 2021; Lynch-Stieglitz et al., 2015; Costa and McManus, 2017
ML1208-27BB	Lynch-Stieglitz et al., 2015; Monteagudo et al., 2021; Lynch-Stieglitz, unpublished
ML1208-28BB	Lynch-Stieglitz et al., 2015; Costa et al., 2016b; Costa and McManus, 2017; Monteagudo et al., 2021; Lynch-Stieglitz, unpublished
ML1208-30BB	Lynch-Stieglitz et al., 2015
ML1208-31BB	Lynch-Stieglitz et al., 2015; Jacobel et al., 2016; Monteagudo et al., 2021; Lynch-Stieglitz, unpublished
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ML1208-34BB	Lynch-Stieglitz et al., 2015
ML1208-35BB	Lynch-Stieglitz et al., 2015

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ML1208-36BB	Costa et al., 2016b; Costa and McManus, 2017; Monteagudo et al., 2021; Lynch-Stieglitz, unpublished
ML1208-37BB	Lynch-Stieglitz et al., 2015; Jacobel et al., 2016; Monteagudo et al., 2021
MR00-K03-PC-01	Harada et al., 2004
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MS21PC	Hennekam et al., 2015
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NEAP-04K	Rickaby and Elderfield, 2005; Hall et al., 2004
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OC205-2-100GGC	Slowey and Curry, 1995; Came et al., 2008
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OCE326-26GGC	Keigwin et al., 2005
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ODP1063	Channell et al., 2012
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ODP769	Linsley, 1996
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OK92_2182	Kaiser, 2002
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Orgon4-KS8	Sirocko, 1989; Sirocko et al., 2000
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PASSAP_PS009PC	Hennekam et al., 2015
PC17	Lee et al., 2001
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PLDS-7G	Keigwin and Lehman, 2015
POS200_10_6-2	Abrantes et al., 2018; Abrantes et al., 2001; Abrantes et al., 1998; Baas et al., 1997; Mienert et al., 1998
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PS1006-1	Grobe and Mackensen, 1992
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PS1481-3	Grobe and Fütterer, 1990
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PS1494-3	Melles, 1991
PS1498-1	Melles, 1991
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PS1506-1	Mackensen et al., 1994
PS1519-12	Horwege/Spielhagen, unpublished
PS1524-1	Köhler, 1991
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PS1535-5	Spielhagen et al., 2004; Nørgaard-Pedersen et al., 2003
PS1535-8	Spielhagen et al., 2004; Nowaczyk et al., 2003
PS1563-2	Grobe, 2002a
PS1564-2	Grobe, 2002b
PS1565-2	Hillenbrand, 1995
PS1576-2	Brehme, 1992
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PS1599-3	Weber, 1992; Weber et al., 1994
PS1606-3	Melles, 1991
PS1607-1	Melles, 1991
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PS1609-3	Melles, 1991
PS1611-3	Melles, 1991
PS1612-1	Melles, 1991
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PS1640-1	Grobe and Mackensen, 1992
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PS1649-2	Ott and Gersonde, 1997b
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PS1654-2	Ott and Gersonde, 1997l; Bianchi and Gersonde, 2004
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PS2498-1	Mackensen et al., 2001; Gersonde et al., 2003; Niebler, 2004d; Niebler, 2004e; Niebler, 2004f
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PS75-059-2	Ullermann et al., 2016; Ronge et al., 2016
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T86-15P	Sarnthein et al., 1994
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TAN0803-09	Maxson et al., 2019; Bostock et al., 2015
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V19-28	Koutavas and Lynch-Stieglitz, 2003
V19-30	Curry et al., 1988
V20-234	Lynch-Stieglitz, unpublished
V21-146	Hovan et al., 1991
V21-29	Koutavas and Lynch-Stieglitz, 2003
V21-30	Koutavas and Lynch-Stieglitz, 2003
V21-40	Koutavas and Lynch-Stieglitz, 2003
V22-108	Charles et al., 1991
V22-174	Shackleton, 1977b
V22-196	Sarnthein et al., 1994
V22-197	Curry et al., 1988
V22-222	Mix et al., 1986
V23-100	Sarnthein et al., 1994
V23-81	Jansen and Veum, 1990; Elliot et al., 1998
V24-109	Shackleton et al., 1992
V24-157	Anderson et al., 1989
V24-161	Anderson et al., 1989
V24-166	Anderson et al., 1989
V24-170	Anderson et al., 1989
V24-184	Anderson et al., 1989
V24-253	Oppo and Horowitz, 2000
V25-21	Curry and Crowley, 1987
V25-59	Curry et al., 1988
V26-175	Matsumoto and Lynch-Stieglitz, 2003
V26-176	Sarnthein et al., 1988; Matsumoto and Lynch-Stieglitz, 2003; CLIMAP Project Members, 2004b
V26-177	Matsumoto and Lynch-Stieglitz, 2003
V27-180	Lynch-Stieglitz, unpublished

<b>Core/Site</b>	<b>References</b>
V28-122	Oppo and Fairbanks, 1987; Broecker et al., 1988a; Broecker et al., 1988b; Schmidt et al., 2004
V28-127	Oppo and Fairbanks, 1990
V28-14	Curry et al., 1988
V28-304	Curry et al., 1988
V28-73	Oppo and Lehman, 1993
V29-135	Sarnthein et al., 1994
V29-140	Lynch-Stieglitz et al., 2006
V29-193	Oppo and Lehman, 1993
V29-198	Oppo and Lehman, 1993
V29-202	Oppo and Lehman, 1993
V29-204	Curry et al., 1999
V29-9	Lynch-Stieglitz, unpublished
V30-40	Oppo and Fairbanks, 1987
V30-49	Curry et al., 1988
V30-5	Matsumoto and Lynch-Stieglitz, 2003
V32-8	Mix et al., 1986
V34-90	Gorbarenko et al., 2002
V34-98	Gorbarenko et al., 2002
V35-5	Oppo and Fairbanks, 1987
Vi-37GC	Keigwin, 1998
VM12-107	Schmidt et al., 2012
VM18-222	Lynch-Stieglitz et al., 1994
VM19-110	Leech et al., 2013
VM24-110	Leech et al., 2013
VM24-150	Leech et al., 2013
VM28-213	Leech et al., 2013
VM28-227	Leech et al., 2013
VM28-229	Leech et al., 2013
VM28-230	Leech et al., 2013
VM28-233	Leech et al., 2013
VM28-234	Leech et al., 2013

Core/Site	References
VM28-235	Leech et al., 2013
VM28-235TW	Leech et al., 2013
VM28-236	Leech et al., 2013
VM28-246	Leech et al., 2013
VM34-2	Leech et al., 2013
VNTR01_10PC	Keigwin and Lehman, 2015
W8402A-14	Jasper et al., 1994
W8709A-1	Lyle et al., 1992
W8709A-13	Lyle et al., 1992; Lund and Mix, 1998
W8709A-8	Lyle et al., 1992; Ortiz et al., 1997
W8709A-8TC	Lyle et al., 1992; Ortiz et al., 1997
WIND-28K	Kiefer et al., 2006; Johnstone et al., 2014
Y71-06-12	Shackleton, 1977a
Y71-09-101	Lyle et al., 2002
Z2108	Nelson et al., 1994
Z2112	Sikes et al., 2016

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