





20

## Abstract

21 Surface incident solar radiation ( $R_s$ ) plays an essential role in climate change on  
22 Earth.  $R_s$  can be directly measured, and it shows substantial variability, i.e., global  
23 dimming and brightening, on decadal scales.  $R_s$  can also be derived from the observed  
24 sunshine duration (SunDu) with reliable accuracy. The SunDu-derived  $R_s$  was used as  
25 a reference to detect and adjust the inhomogeneity in the observed  $R_s$ . However, both  
26 the observed  $R_s$  and SunDu-derived  $R_s$  may have inhomogeneity. In Japan, SunDu has  
27 been measured since 1890, and  $R_s$  has been measured since 1961 at ~100 stations. In  
28 this study, the observed  $R_s$  and SunDu-derived  $R_s$  were first checked for inhomogeneity  
29 with a statistical software RHtest. If confirmed by the metadata of these observations,  
30 the detected inhomogeneity was adjusted based on the RHtest-quantile matching  
31 method. Second, the two homogenized time series were compared to detect further  
32 possible inhomogeneity. If confirmed by the independent ground-based observations of  
33 cloud cover fraction, the detected inhomogeneity was adjusted based on the reference  
34 dataset. As a result, a sharp decrease in the observed  $R_s$  from 1961 to 1975 caused by  
35 instrument displacement was detected and adjusted. Similarly, a gradual decline in  
36 SunDu-derived  $R_s$  due to steady instrument replacement from 1985 to 1990 was  
37 detected and adjusted. After homogenization, the two estimates agree well.  $R_s$  was  
38 found to have increased at a rate of  $0.9 \text{ W m}^{-2}$  per decade ( $p < 0.01$ ) from 1961 to 2015  
39 based on the homogenized SunDu-derived  $R_s$ , which was enhanced by a positive  
40 aerosol-related radiative effect ( $2.2 \text{ W m}^{-2}$  per decade) and diminished by a negative



41 cloud cover radiative effect ( $-1.4 \text{ W m}^{-2}$  per decade). The brightening over Japan was  
42 the strongest in spring, likely due to a significant decline in aerosol transported from  
43 Asian dust storms. The observed raw  $R_s$  data and their homogenized time series used in  
44 this study are available at <https://doi.org/10.11888/Meteoro.tpdc.271524> (Ma et al.,  
45 2021).

46

47 **Key Points:**

48 (1) Surface incident solar radiation ( $R_s$ ) and sunshine duration over Japan were  
49 homogenized.

50 (2) Homogenized century-long  $R_s$  data over Japan were produced, and shows that  $R_s$   
51 increased at a rate of  $\sim 1 \text{ W m}^{-2}$  per decade from 1961 to 2015.

52 (3) Cloud cover modulates  $R_s$  variation at monthly and interannual time scales, while  
53 aerosols dominate the decadal variation in  $R_s$ .



## 54 **1. Introduction**

55       Surface incident solar radiation ( $R_s$ ) plays a vital role in atmospheric circulation,  
56 hydrologic cycling and ecological equilibrium; therefore, its decrease and increase  
57 termed as global dimming and brightening (Wild et al., 2005; Shi et al., 2008), have  
58 received widespread interest from the public and scientific community (Allen et al.,  
59 2013; Xia, 2010; Wang et al., 2013; Tanaka et al., 2016; Ohmura, 2009; Stanhill and  
60 Cohen, 2005).  $R_s$  can be measured by either a single pyranometer or the summation of  
61 diffuse and direct components. The measurement of  $R_s$ , which started in 1961 in Japan,  
62 has a long history (Tanaka et al., 2016), and a data record more than half a century-long  
63 has been accumulated. The dataset has been widely used to study decadal variability  
64 (Wild et al., 2005; Stanhill and Cohen, 2008) and to evaluate model simulations (Allen  
65 et al., 2013; Dwyer et al., 2010).

66       The Eppley and Robitzsch pyranometers used to measure  $R_s$  over Japan were  
67 replaced by the Moll-Gorczynski thermopile pyranometers in the early 1970s (Tanaka  
68 et al., 2016). Instrument replacements introduced substantial inhomogeneity into the  
69 time series of observed  $R_s$  over China during the period of 1990-1993 (Shi et al., 2008;  
70 Wang et al., 2015). Instrument changes from the Robitzsch pyranograph to the Kipp &  
71 Zonen CM11 pyranometer before 1980 caused no clear dimming in Italy (Manara et al.,  
72 2016). However, the possible homogeneity of the observed  $R_s$  over Japan has not been  
73 well quantified, and most existing studies directly used raw  $R_s$  data (Wild et al., 2005;



74 Tanaka et al., 2016; Tsutsumi and Murakami, 2012; Allen et al., 2013; Wild and  
75 Schmucki, 2011; Kudo et al., 2012; Ohmura, 2009). Some studies have had to abandon  
76 data from the early years and focused on only  $R_s$  data collected after 1975 (Tsutsumi  
77 and Murakami, 2012; Dwyer et al., 2010). Therefore, the observed decadal variability  
78 in  $R_s$  over Japan is questionable, especially for the 1961-1975 time period.

79 Homogenizing the observed  $R_s$  has been attempted in China (Wang et al., 2015;  
80 Tang et al., 2011; Yang et al., 2018), Italy (Manara et al., 2016), Spain (Sanchez-  
81 Lorenzo et al., 2013) and Europe (Sanchez-Lorenzo et al., 2015). It is essential to find  
82 a homogeneous reference station to compare with the possible inhomogeneous station  
83 to test and adjust the inhomogeneity in the observed time series, as done for the  
84 homogenization of air temperature (Du et al., 2020). However, this process is difficult  
85 for  $R_s$  because the instrument replacement of  $R_s$  generally occurs nearly simultaneously  
86 throughout a country. Therefore, the sunshine duration (SunDu)-derived  $R_s$  (Yang et al.,  
87 2006) has been used as a homogeneous reference dataset to detect and adjust the  
88 inhomogeneity of  $R_s$  in China (Wang et al., 2015).

89 The SunDu records the hours of surface direct solar radiation exceeding  $120 \text{ W m}^{-2}$   
90 <sup>2</sup> and provides an alternative way to estimate  $R_s$  (Yang et al., 2006; Stanhill and Cohen,  
91 2008). It has been revealed that the Jordan recorder is 10% more sensitive than the  
92 Campbell-Stokes recorder for SunDu measurements (Noguchi, 1981). The  
93 homogenization of SunDu has been carried out in Iberian Peninsula (Sanchez-Lorenzo  
94 et al., 2007), Switzerland (Sanchez-Lorenzo and Wild, 2012), and Italy (Manara et al.,



95 2015). In Japan, the Jordan recorders used to measure SunDu were replaced by EKO  
96 rotating mirror recorders in approximately 1986 (Inoue and Matsumoto, 2003; Stanhill  
97 and Cohen, 2008). Therefore, SunDu observations over Japan themselves may have  
98 inhomogeneity issues.

99 The RHtest-quantile matching (QM) method (Wang, 2008b; Vincent et al., 2012),  
100 which first detects the changepoints in a series and then tunes the inhomogeneous data  
101 segments to be consistent with other segments in empirical distributions, has been  
102 widely used for homogenizing climate variables (Dai et al., 2011; Wang et al., 2010).

103 The main objective of this study is to detect and adjust the inhomogeneity in  $R_s$   
104 estimates over Japan. The metadata were first extracted from website information and  
105 related records at each site. The SunDu observations were converted into  $R_s$ . The  
106 RHtest-QM method was applied to homogenize the observed  $R_s$  and SunDu-derived  $R_s$ ,  
107 and finally, the homogenized long-term  $R_s$  data were derived over Japan. Furthermore,  
108 the impacts of cloud cover and aerosols on  $R_s$  variation over Japan in recent decades  
109 were explored.

## 110 **2. Data and methods**

### 111 **2.1 Surface incident solar radiation and sunshine duration**

112 The monthly observed  $R_s$  at 105 stations and SunDu at 156 stations were  
113 downloaded from the Japanese Meteorology Agency (JMA) website (see Table S1 and  
114 Figure 1).  $R_s$  records were available from 1961. During the 1960s, two  $R_s$  measurements



115 were conducted in parallel by both Eppley and Robitzsch pyranometers. In the early  
116 1970s (see Figure 2 and Table S2), these instruments were replaced by Moll-Gorczyński  
117 thermopile pyranometers. This replacement occurred at approximately 12% of  $R_s$   
118 stations in 1971, followed by ~24%, 26%, 4% and 32% in the next four years, which  
119 may have caused severe data discontinuity problems (Tanaka et al., 2016).

120 SunDu has been routinely measured since 1890. Jordan recorders were replaced  
121 by EKO rotating mirror recorders at nearly 50% of SunDu stations in 1986. Before  
122 1990, nearly all of the SunDu stations used new instruments for observations. Less than  
123 5% of SunDu stations before 1985 and more than 10% of SunDu stations after 2000  
124 were moved away from the original sites (see Figure 2 and Table S2) (Stanhill and  
125 Cohen, 2008).

126 In this study, SunDu was used to derive  $R_s$  based on the following equation (Yang  
127 et al., 2006):

$$128 \quad R_s / R_c = a_0 + a_1 \cdot n / N + a_2 \cdot (n / N)^2 \quad (1)$$

129 where  $n$  is sunshine duration hours;  $N$  is the maximum possible sunshine duration;  $R_c$   
130 is surface solar radiation under clear skies; and  $a_0$ ,  $a_1$  and  $a_2$  are coefficients. This  
131 method was recommended in many studies (Wang et al., 2015; Tang et al., 2011).

## 132 2.2. Homogenization method

133 Both  $R_s$  and SunDu measurements over Japan suffer inhomogeneity problems,  
134 which require rigorous data homogenization. RHtest  
135 (<http://etccdi.pacificclimate.org/software.shtml>) is a widely used method to detect and



136 adjust multiple changepoints in a data series (Wang, 2008a). Two algorithms were  
137 provided to detect changepoints based on the penalized maximal T test and the  
138 penalized maximal F test (Wang, 2008b). As the discontinuity dates were recorded on  
139 the JMA websites, we artificially treat these observations on those dates as changepoints.  
140 To diminish all significant artificial shifts caused by the changepoints, Quantile-  
141 Matching (QM) adjustments in the RHtest (Vincent et al., 2012) were performed to  
142 adjust the series so that the empirical distributions of all segments of the detrended base  
143 series agree with each other. The corrected values are all based on the empirical  
144 frequency of the datum to be adjusted.

145 Another independent homogenization method proposed by (Katsuyama, 1987),  
146 which was developed due to the replacement of the Jordan recorders with EKO rotating  
147 mirror recorder during the late 1980s, is denoted as follows:

$$148 \quad S_R = 0.8 S_J (S_J < 2.5 \text{ h/day}) \quad (2)$$

$$149 \quad S_R = S_J - 0.5 \text{ h/day} (S_J \geq 2.5 \text{ h/day}) \quad (3)$$

150 where  $S_J$  is the daily SunDu observed by the Jordan recorders before replacement; and  
151  $S_R$  is the daily SunDu adjusted to be consistent with the values observed with the EKO  
152 rotating mirror recorders. The homogenization methods were compared in this study  
153 and yielded nearly the same SunDu-derived  $R_s$  variation, as shown in Figure 3.

### 154 **2.3 Clouds**

155 Clouds play an important role in  $R_s$  variation (Norris and Wild, 2009). Monthly  
156 cloud cover observations at 155 stations were also available on the JMA website. The



157 observation time for cloud amount has been 08:00-19:00 since 1981 at 10% of cloud  
158 amount stations and 08:30-17:00 from 1990 to 1995 at another 10% of cloud amount  
159 stations (see Figure 2 and Table S2). However, the difference between annual raw and  
160 homogenized cloud data is trivial, as cloud data are relatively homogeneous in space  
161 compared with  $R_s$  and SunDu observations. A site observation of cloud amount can  
162 represent the value over a large spatial scale, likely leading to few inhomogeneity issues  
163 for cloud data. The Clouds and the Earth's Radiant Energy System (CERES) provides  
164 surface incident solar radiation (Ma et al., 2015) primarily based on the Moderate  
165 Resolution Imaging Spectroradiometer (MODIS) cloud and aerosol products (Kato et  
166 al., 2012).

167 To explore the impact of the cloud cover anomaly on the  $R_s$  variation, the cloud  
168 cover radiative effect (CCRE), defined as the change in  $R_s$  produced by a change in  
169 cloud cover, was proposed by (Norris and Wild, 2009):

$$170 \quad CCRE'(g, y, m) = CC'(g, y, m) \times CRE(g, m) / \overline{CC}(g, m) \quad (4)$$

171 where  $g$  is the grid,  $y$  is the year,  $m$  is the month,  $CCRE'$  is the cloud cover radiative  
172 effect anomaly,  $CC'$  is the cloud cover anomaly,  $\overline{CC}$  is the long-term mean cloud  
173 cover and  $CRE$  is the cloud radiative effect calculated by the  $R_s$  difference under all  
174 sky and clear sky conditions.

175 The residual radiative effect was determined by removing the CCRE anomalies  
176 from the  $R_s$  anomalies. It is noted that a part of the cloud albedo radiative effect  
177 proportional to the cloud amount was contained in the CCRE, as a large cloud amount



178 tends to yield enhanced cloud albedo, whereas another part of the cloud albedo radiative  
179 effect due to the aerosol first indirect effect (more aerosols facilitating more cloud  
180 condensation nuclei may enhance cloud albedo) may be included in the residual  
181 radiative effect, which mainly contains the aerosol radiative effect. In this study, long-  
182 term observations of cloud amount and monthly cloud radiative effect (CRE) data in  
183 the CERES EBAF edition were used following Equation (4) to distinguish the cloud  
184 cover radiative effect from  $R_s$  variation.

### 185 **3. Results**

#### 186 **3.1 Homogenization of observed $R_s$ and sunshine duration derived $R_s$**

187 In this study, monthly values were converted into annual values for calculation.  
188 If there are missing values in any month in a specific year, the annual value for that  
189 year is set to a missing value. Both  $R_s$  and SunDu records are available at 105 stations.  
190 Figure 4 shows the comparisons between raw data and homogenized data. After QM  
191 adjustments, the correlation coefficients between the annual observed  $R_s$  and annual  
192 SunDu-derived  $R_s$  are significant with a 90% confidence interval at 75 stations. The  
193 correlation coefficients were improved at 54 of 75 stations after homogenization,  
194 including 31 stations that had improvements greater than 0.2. Among the 54 stations,  
195 there were 41 stations (marked with red in Table S1) at which the correlation  
196 coefficients were greater than 0.5, and the biases and the root mean square errors  
197 generally decrease after homogenization.



198           Figure 5 shows the time series of surface incident solar radiation ( $R_s$  and SunDu-  
199    derived  $R_s$ ) at the HAMADA site (WMO-ID: 47755, Lat: 34.9, Lon: 132.07) before  
200    and after homogenization, which highlights the necessity and feasibility of the RHtest-  
201    QM method. The SunDu-derived  $R_s$  variation over Japan during recent decades  
202    inferred from these “perfect” data at 41 sites (Figure 6) was nearly identical to that  
203    from all available data at 156 sites (as shown in Table 1 and Figure 7).

204           The cloud amount in CERES agrees well with the observations, and the annual  
205    CRE in CERES is well correlated with the annual cloud amount in Figure 8. The  
206    regional average cloud amount over Japan in Figure 8 (blue line) increases at a rate of  
207    0.7% per decade from 1960 to 2015, which is consistent with the results (Figure 4) in  
208    (Tsutsumi and Murakami, 2012).

### 209    **3.2 Uncertainties in $R_s$ observations**

210           Figure 7 displays the change in  $R_s$  during the last 5 decades, while Figure 8 shows  
211    the variation in observed clouds over Japan. The sharp decrease in  $R_s$  in 1963 was  
212    attributed to the volcanic eruption of Agung in Indonesia in the same year (Witham,  
213    2005). The sharp decreases in  $R_s$  in 1991 and 1993 are due to the combined effect of  
214    the volcanic eruption of Mount Pinatubo in the Philippines in 1991 (Robock, 2000) and  
215    the simultaneous significant increases in clouds (shown in Figure 8) (Tsutsumi and  
216    Murakami, 2012). The volcanic eruption of El Chichón in Mexico in 1982 exerted little  
217    impact on the decline in  $R_s$  and may have been compensated by the decrease in clouds,  
218    as shown in Figure 8. The pronounced  $R_s$  decline in 1980 coincides with the significant



219 increase in clouds, while the lightening of  $R_s$  in 1978 and 1994 encounters abrupt  
220 decreases in cloud covers.

221 As shown in Figure 7,  $R_s$  observations change little after homogenization  
222 (comparison between the light blue and dark blue lines). However, the SunDu-derived  
223  $R_s$  series are smoother after adjustment by the QM method, as the sharp decrease from  
224 1983 to 1993 caused by the replacement of sunshine duration instruments (Jordan  
225 recorders were replaced with EKO rotating mirror recorders) (Stanhill and Cohen, 2008)  
226 was repaired (comparison between the light red line and dark red lines). Despite the  
227 identical increase in  $R_s$  via both the homogenized direct measurements of  $R_s$  and the  
228 homogenized SunDu-derived  $R_s$  during the 1995-2014 period, their variations in  $R_s$   
229 from 1961 to 1994 are different (dark red line and dark blue line).

230 Large discrepancies in  $R_s$  variation were found during the time period of 1961-  
231 1970, although homogenizations were performed on the direct measurements of  $R_s$  and  
232 SunDu-derived  $R_s$  (dark blue line and dark red line in Figure 7). Existing study noted  
233 the inaccurate instruments used at the beginning of operation of the  $R_s$  observation  
234 network in approximately 1961, and the parallel use of two different types of  
235 instruments during the 1960s may result in the large variability in observed  $R_s$  (Tanaka  
236 et al., 2016). At this time, the clouds fluctuated gently, as shown in Figure 8, and the  
237 change in volcanic aerosols from 1965 to 1966 was nearly the same as that from 1962  
238 to 1963 (Table 2 in (Sato et al., 1993)), so the sudden decline in the direct observations  
239 of  $R_s$  from 1965 to 1966, which was twice as large as that from 1962 to 1963, is



240 suspicious. It is inferred that anthropogenic aerosols play a subtle role in the significant  
241 reduction in  $R_s$ , as this type of phenomenon is common for both polluted and pristine  
242 stations in Japan (Figure 22 in (Tanaka et al., 2016)).

243 Figure 9 shows the correlation coefficients between homogenized  $R_s$  (observed  
244 and SunDu-derived) and cloud amount. In general, the observed  $R_s$  (-0.45) is less  
245 correlated than the SunDu-derived  $R_s$  (-0.67), particularly from 1961 to 1970, -0.21  
246 compared with -0.64. This in turn supports the reliability of homogenized SunDu-  
247 derived  $R_s$ , especially during the time period of 1961-1970. The misleading  $R_s$  variation  
248 was modified by the RHtest method again using homogenized SunDu-derived  $R_s$  as  
249 reference data from 1961 to 1970 as shown in Figure 10.

250 General decreases in stratospheric aerosol optical depth (AOD) were reported in  
251 (Sato et al., 1993) from 1965 to 1980, and clouds fluctuated slightly, as shown in Figure  
252 8; both of these factors contributed to a brightening of  $R_s$ . This is in agreement with the  
253 SunDu-derived  $R_s$  and contrasts with the direct measurements of  $R_s$ .

254 During the 1985-1990 period, clouds varied slightly, as shown in Figure 8, and the  
255 observed atmospheric transmission under cloud-free conditions increased (Wild et al.,  
256 2005), which suggests that the large declines in directly observed  $R_s$  and SunDu-derived  
257  $R_s$  are defective and reinforce the reliability of the adjusted SunDu-derived  $R_s$  (dark red  
258 line in Figure 7).

259 From the above analysis, it can be inferred that fewer uncertainties exist in  
260 homogenized SunDu-derived  $R_s$ , which was confirmed by another work that utilized a



261 different data adjusted method (Stanhill and Cohen, 2008).

### 262 **3.2 Trends of $R_s$ over Japan**

263 The trends of  $R_s$  during specific time periods for different types of datasets are  
264 listed in Table 1. Direct measurements of  $R_s$  and SunDu-derived  $R_s$  from 41 selected  
265 stations and all available stations reveal similar variations in  $R_s$  over Japan, which  
266 demonstrates that the sample number has a subtle impact on the estimation of global  
267 brightening and dimming over Japan.

268 Major differences were found in the time periods of 1961-1980, ranging from -  
269 11.2 (-12.0) to -8.4 (-4.8)  $\text{W m}^{-2}$  per decade before and after  $R_s$  homogenizations for all  
270 available stations (41 selected stations) over Japan. In addition, significant repairs  
271 occurred during the 1981-1995 period, ranging from -10.6 (-11.3) to -1.2 (-1.3)  $\text{W m}^{-2}$   
272 per decade before and after SunDu-derived  $R_s$  homogenizations for all available stations  
273 (41 selected stations) over Japan. Both corrections were mainly attributed to the  
274 homogenization of corrupted raw data caused by the replacement of instruments for  $R_s$   
275 and SunDu measurements. After careful checking and adjustment of the SunDu-derived  
276  $R_s$  series, the decadal variation in  $R_s$  over Japan, which was totally different from former  
277 studies (Wild et al., 2005; Norris and Wild, 2009), was remedied.

278 The combined effects of clouds and aerosols on  $R_s$  make the global dimming and  
279 brightening complicated. The CCRE can explain 70% of global brightening from 1961  
280 to 2014 at monthly and interannual time scales, while the residual radiative effect  
281 dominates the decadal variation in  $R_s$ , as shown in Figure 11 and Table 1, which is in



282 agreement with Wang et al. (2012). Homogenized SunDu-derived  $R_s$  show an increase  
283 of  $1.6 \text{ W m}^{-2}$  per decade from 1961 to 1980; however, persistent increase in cloud  
284 amount yields a CCRE decrease of  $1.1 \text{ W m}^{-2}$  per decade. The residual radiative effect  
285 accounts for an increase of  $2.4 \text{ W m}^{-2}$  per decade for this time period. The cloud  
286 radiative effect ( $-1.4 \text{ W m}^{-2}$  per decade) modulates  $R_s$  variation of  $-1.2 \text{ W m}^{-2}$  per decade  
287 for the 1981-1995 period, while the residual radiative effect ( $1.2 \text{ W m}^{-2}$  per decade)  
288 dominates  $R_s$  variation of  $1.4 \text{ W m}^{-2}$  per decade from 1996 to 2014.

289 Homogenized SunDu-derived  $R_s$  shows a slight increase of  $0.9 \text{ W m}^{-2}$  per decade  
290 from 1961 to 2014 with a 90% confidence interval. However, the CCRE accounts for a  
291 decreased  $R_s$  of  $1.4 \text{ W m}^{-2}$  per decade, which implies that cloud cover changes are not  
292 the primary driving forces for the  $R_s$  trend over Japan. Meanwhile, the residual radiative  
293 effect exhibits an increase of  $2.2 \text{ W m}^{-2}$  per decade, which surpasses the negative CCRE.

294 Several studies demonstrate a generally cleaner sky over Japan from the 1960s to  
295 the 2000s (except for the years impacted by volcanic eruptions) based on atmospheric  
296 transparency and aerosol optical properties (Wild et al., 2005; Kudo et al., 2012), which  
297 supports the dominant role of aerosols in  $R_s$  brightening over Japan, as revealed by the  
298 residual radiative effect here. Furthermore, the residual radiative effect in this study is  
299 stronger than that in Norris and Wild (2009), as raw data were remedied and more  
300 accurate satellite data from CERES were adopted to quantify the radiative effect.  
301 Tsutsumi and Murakami (2012) demonstrated that cloud amount categories exert an  
302 important effect on  $R_s$  variation.  $R_s$  enhancement by the increased appearance of large



303 cloud amounts is superior to  $R_s$  decline by the decreased appearance of small cloud  
304 amounts during 1961-2014, which yields increased  $R_s$  with increasing total cloud  
305 amount. They also pointed out that the decrease in cloud optical thickness due to the  
306 large emissions of  $\text{SO}_2$  and black carbon from East Asia through the aerosol semi-direct  
307 effect (absorption of more energy by aerosols results in the evaporation or suppression  
308 of clouds) may have facilitated the increased  $R_s$  over Japan.

309 The decrease in spring dust storms in March-May during the last 5 decades from  
310 China (Qian et al., 2002; Zhu et al., 2008), which may travel to neighbouring  
311 countries(Uno et al., 2008; Choi et al., 2001), could also have triggered the increase in  
312  $R_s$  over Japan. The  $R_s$  variation and radiative effect in different seasons are categorized  
313 in Figure 12 and Table 2, in which an increasing trend of  $1.5 \text{ W m}^{-2}$  per decade in the  
314 homogenized SunDu-derived  $R_s$  prevails in spring for the whole time period, dominated  
315 by a dramatic increase of  $2.8 \text{ W m}^{-2}$  per decade in the residual effect and even larger  
316 increase during 1961-1980 ( $3.1 \text{ W m}^{-2}$  per decade) and 1996-2014 ( $3.4 \text{ W m}^{-2}$  per  
317 decade).

#### 318 **4. Data availability**

319 Monthly observed surface incident solar radiation, sunshine duration and cloud amount  
320 data were provided by Japan Meteorological Agency  
321 (<https://www.data.jma.go.jp/obd/stats/data/en/smp/index.html>), and monthly cloud  
322 radiative effect (CRE) data were derived from Clouds and the Earth's Radiant Energy  
323 System for CERES EBAF data ([https://ceres.larc.nasa.gov/order\\_data.php](https://ceres.larc.nasa.gov/order_data.php)). The



324 homogenized observed  $R_s$  and SunDu-derived  $R_s$  used in this study are available at  
325 <https://doi.org/10.11888/Meteoro.tpdc.271524> (Ma et al., 2021).

## 326 **5. Conclusions**

327 Observational data themselves have inherent problems caused by measurement  
328 method, instrument replacement and site relocation. Therefore, precautions should be  
329 taken when using these data for trend analysis or as validation data. In this study, the  
330 RHtest-QM method was introduced to homogenize the direct measurements of  $R_s$  and  
331 SunDu-derived  $R_s$  over Japan using the information in metadata as changepoints.  
332 Inhomogeneities in the homogenized raw  $R_s$  was further checked by exploring the  
333 relationship with the ground-based cloud amount and tuned again using homogenized  
334 SunDu-derived  $R_s$  as the reference data. The global dimming and brightening over  
335 Japan were revisited based on the homogenized SunDu-derived  $R_s$ , which diminished  
336 the effect of nonclimate signals in the raw observations.

337 Japan experienced a sudden decline in  $R_s$  in 1963, a global brightening of  $4.8 \text{ W}$   
338  $\text{m}^{-2}$  per decade ( $P < 0.01$ ) from 1963 to 1977, a rapid increase in 1978, a sudden decrease  
339 in 1980, a global dimming of  $5.1 \text{ W m}^{-2}$  per decade ( $P < 0.10$ ) from 1981 to 1993, a  
340 pronounced increase in 1994, and a nearly  $1 \text{ W m}^{-2}$  per decade increase from 1995 to  
341 2014. For the last 5 decades, a slight global brightening of  $1 \text{ W m}^{-2}$  per decade (with a  
342 99% confidence interval) was inferred from the homogenized SunDu-derived  $R_s$ .  
343 Global brightening since 1961 over Japan is consistent with that in (Stanhill and Cohen,



344 2008), except that the magnitude is not as large.

345       Clouds and aerosols are the two major factors that mediate the transformation of  
346  $R_s$ . The brightening in Japan for 1961-1980 was the combined effect of cloud cover  
347 (negative effect) and aerosols (positive effect). The dimming for 1981-1995 was  
348 governed by reduced cloud amounts, while the increase in  $R_s$  for 1996-2014 was  
349 controlled by decreased aerosols. These results are different from those in (Norris and  
350 Wild, 2009), as homogenization was performed on the raw data and more accurate  
351 cloud radiative effect data series from CERES were utilized in our study. During the  
352 entire period of 1961-2014, cloud amounts dominated seasonal and interannual  $R_s$   
353 variations, while aerosols (including aerosol-cloud interactions) drove decadal  $R_s$   
354 variations over Japan, noted by other studies, in response to generally cleaner skies and  
355 a reduction in spring Asian dust storms (Wang et al., 2012; Kudo et al., 2012).  
356



357 **Author contributions**

358 QM and KW designed the research and wrote the paper. LS collected the raw data. YH  
359 homogenized the raw data. QW provided the technical support. YZ and HL checked the  
360 data.

361

362 **Competing interests**

363 The authors declare that they have no conflict of interest.

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### Acknowledgements

368 This study is funded by the National Key R&D Program of China (2017YFA0603601),  
369 the National Science Foundation of China (41930970), and Project Supported by State  
370 Key Laboratory of Earth Surface Processes and Resource Ecology (2017-KF-03). We  
371 thank many institutions for sharing their data: Japan Meteorological Agency for  
372 observation data over Japan  
373 (<https://www.data.jma.go.jp/obd/stats/data/en/smp/index.html>); Clouds and the Earth's  
374 Radiant Energy System for CERES EBAF data  
375 ([https://ceres.larc.nasa.gov/order\\_data.php](https://ceres.larc.nasa.gov/order_data.php)). We thank the Expert Team on Climate  
376 Change Detection and Indices (ETCCDI) for providing the RHtestV4 homogenization  
377 package (<http://etccdi.pacificclimate.org/software.shtml>).

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519 Table 1. Trends of Surface Incident Solar Radiation ( $R_s$ ) in Japan during Specific Time  
 520 Periods for Different Types of Datasets<sup>a</sup>. Unit:  $W\ m^{-2}$  per decade  
 521

Case <sup>b</sup>	Datasets <sup>c</sup>	1961-1980	1981-1995	1996-2014	1961-2014
Selected 41 Stations	OBS	-12.0**	-2.1	2.4	-0.3
	OBS_HM	-4.8*	-2.1	2.4	1.5**
	OBS_2HM	-0.8*	-2.1	2.4*	0.9**
	SunDu-derived	1.4	-11.3**	1.4	-2.1**
	SunDu-derived_HM	1.4	-1.3*	1.5	0.9**
All Stations	OBS	-11.2**	-1.3	2.2	0.2
	OBS_HM	-8.4**	-1.3	2.2	0.8
	OBS_2HM	0.7	-1.3	2.2	1.6**
	SunDu-derived	2.3*	-10.6**	1.2	-1.9**
	SunDu-derived_HM	1.6	-1.2	1.4	0.9*
Radiative Effect	CCRE series	-1.1	-1.4	-0.0	-1.4**
	Residual series	2.4**	-0.1	1.2*	2.2**

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523

524

525 <sup>a</sup>The trend calculations were based on the linear regression method. Values with two  
 526 asterisks (\*\*) imply  $p < 0.01$ , and those with one asterisk (\*) imply  $0.01 < p < 0.1$ .

527 <sup>b</sup> $R_s$  trends were calculated by different numbers of observations, including all stations  
 528 that are available on the JMA website and 41 stations (marked with red in Table S1,  
 529 detailed in Section 3.1) that are significantly improved after homogenization. This  
 530 implies that the sample number has a subtle impact on the trend calculation over Japan.

531 Radiative effects from clouds and aerosols were also explored.

532 <sup>c</sup>Trend calculations were based on the raw measurements of surface incident solar  
 533 radiation (OBS), their homogenized series (OBS\_HM), derived incident solar radiation  
 534 from sunshine duration hours (SunDu-derived) and their homogenized series (SunDu-  
 535 derived\_HM). OBS\_HM from 1961 to 1970 was further homogenized by using SunDu-  
 536 derived\_HM as reference data, termed OBS\_2HM. It is found that homogenized



537 SunDu-derived  $R_s$  have the lowest uncertainties among these five datasets in Section  
538 3.1. The cloud cover radiative effect (CCRE) was denoted as the change in  $R_s$  produced  
539 by a change in cloud cover, and the CCRE calculations were performed following  
540 Equation (4) by observed cloud amounts and the cloud radiative effect (CRE) from  
541 CERES satellite retrieval. Residual effect series were obtained by removing the CCRE  
542 from homogenized SunDu-derived  $R_s$  anomalies.  
543



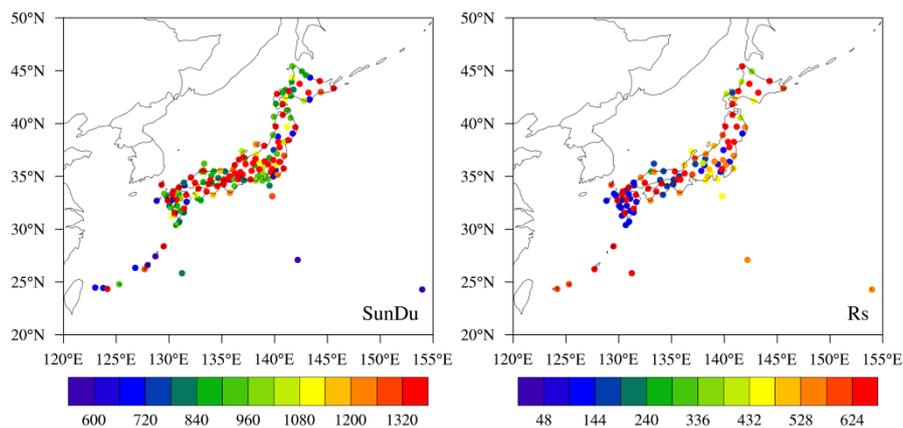
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545 Table 2. Trends of Surface Incident Solar Radiation ( $R_s$ ) in Japan during Specific Time

546 Periods for Different Types of Datasets for All Seasons. Unit:  $\text{W m}^{-2}$  per decade

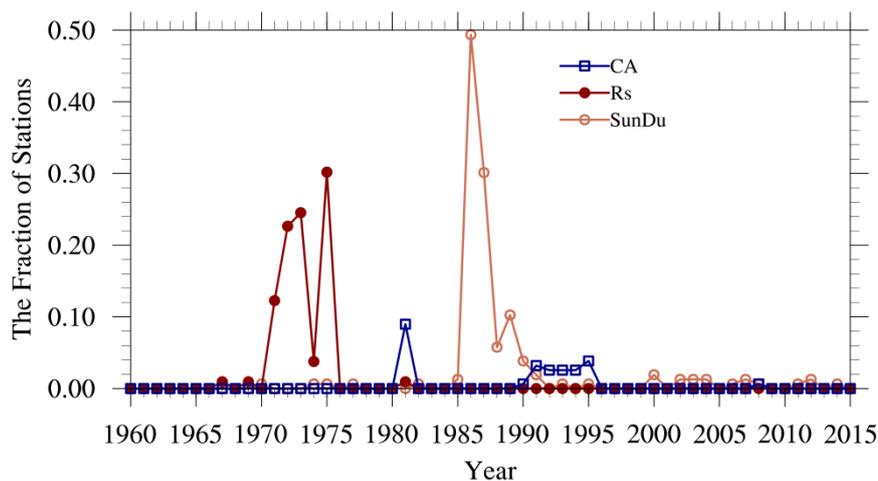
Season	Datasets	1961-1980	1981-1995	1996-2014	1961-2014
Spring	SunDu-derived_HM	3.1	-1.5	3.4*	1.5
	CCRE series	-0.7	-1.6	-1.6	-0.9
	Residual series	4.9**	-0.5**	2.2**	2.8*
Summer	SunDu-derived_HM	1.4	-3.4	0.6	0.4
	CCRE series	-1.9	-2.1	-4.4**	-2.7
	Residual series	2.0**	-1.8	1.5**	2.8
Autumn	SunDu-derived_HM	0.6	1.5	3.3**	1.0*
	CCRE series	-1.3**	1.6	1.6	-0.9
	Residual series	1.8**	0.8**	2.1**	2.0*
Winter	SunDu-derived_HM	0.6	-1.5	-1.6	0.5
	CCRE series	-0.6	-3.3	-0.6	-0.7
	Residual series	1.1**	0.9**	-0.9**	1.2**

547



548

549 Figure 1. The spatial distribution of stations over Japan with observed sunshine duration  
550 (SunDu, 156 stations) and surface incident solar radiation ( $R_s$ , 105 stations) data. The  
551 colours indicate the data length of the SunDu records from 1890 to 2015 and  $R_s$  records  
552 from 1961 to 2015. Unit: month.

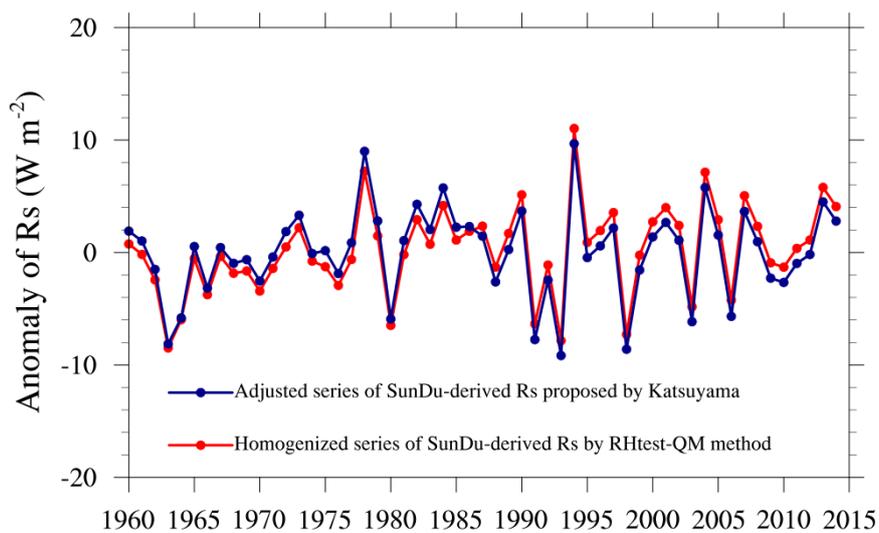


553

554 Figure 2. The fraction of stations that suffer from data inhomogeneity due to site  
555 relocation, change of instruments and measurement method for sunshine duration  
556 (SunDu) records, cloud amount (CA) records and surface incident solar radiation ( $R_s$ )  
557 records. In total, there were 156 stations with SunDu records, 105 of which had  $R_s$   
558 records and 155 of which had CA records. The inhomogeneity information shown here  
559 was derived from metadata from  
560 <https://www.data.jma.go.jp/obd/stats/data/en/smp/index.html>, and was used as primary  
561 information to perform the inhomogeneity adjustment in the RHtest method detailed in  
562 Section 2.2.

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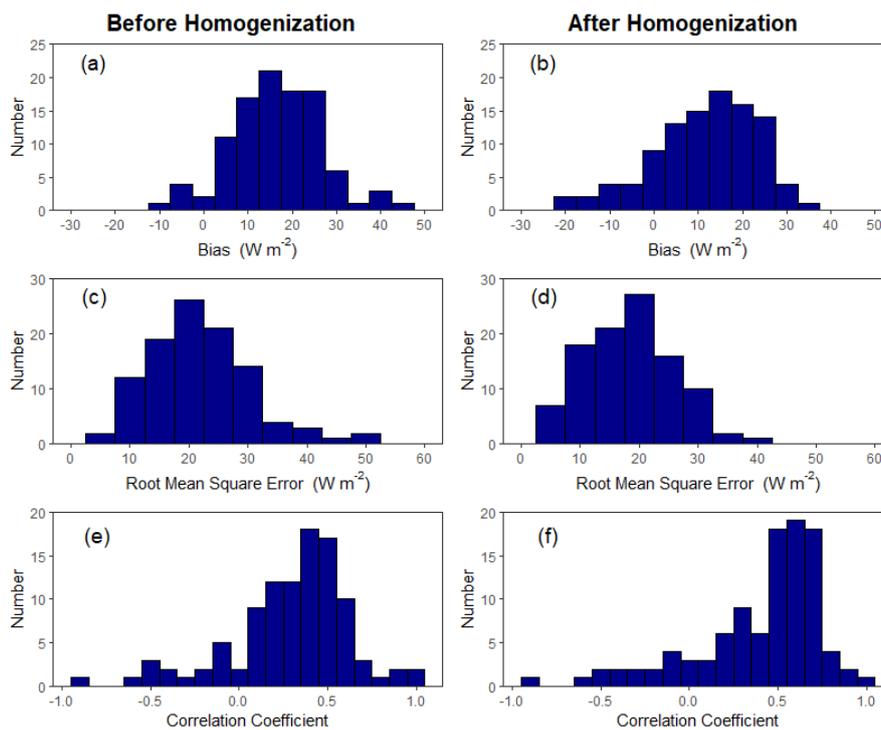


565

566 Figure 3. The anomalies of surface incident solar radiation ( $R_s$ ) derived from  
567 homogenized sunshine duration (SunDu) data (red line) by the RHtest-QM method  
568 and other independent data (blue line) adjusted by the method in (Katsuyama, 1987).  
569 Both of the homogenized datasets yield nearly the same  $R_s$  variation.



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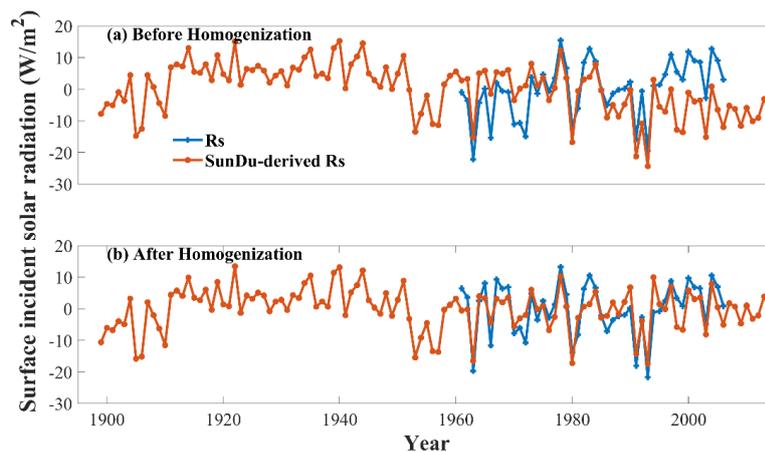
572 Figure 4. Histograms of bias, root mean square error

573 between SunDu-derived surface incident solar radiation ( $R_s$ ) and observed  $R_s$  before

574 (a, c, e) and after (b, d, f) homogenization. Their differences decrease after

575 homogenization.

576

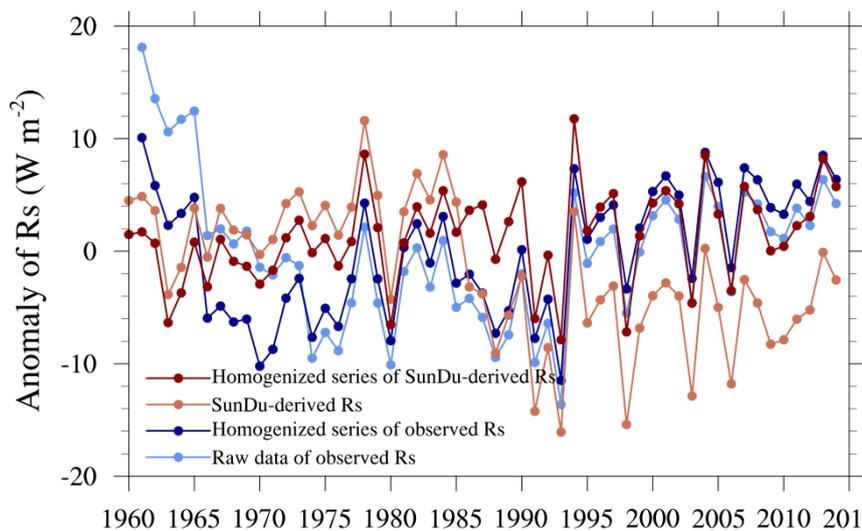


577

578 Figure 5. Time series of annual anomalies of observed surface incident solar radiation  
579 ( $R_s$ ) and SunDu-derived  $R_s$  at HAMADA site (WMO-ID: 47755, Lat:  $34.9^\circ$  , Lon:  
580  $132.07$ ) before and after homogenization.

581

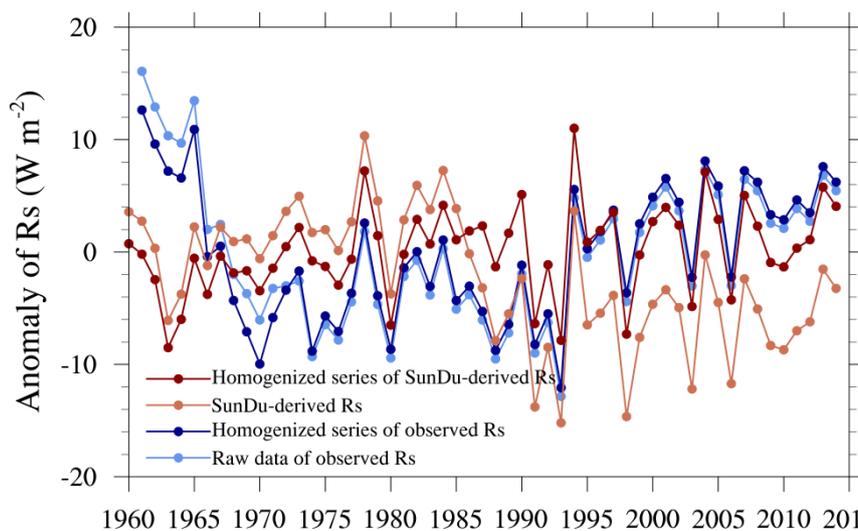
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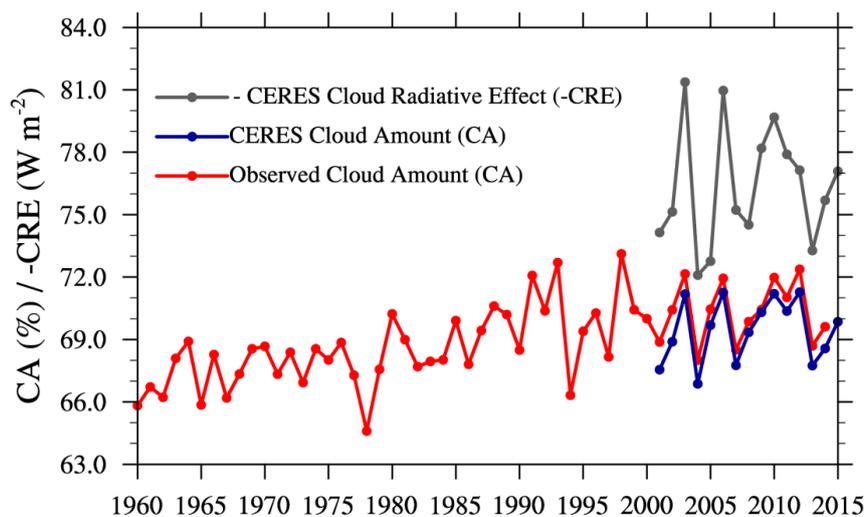
584 Figure 6. Time series of annual anomalies of surface incident solar radiation ( $R_s$ ) based  
585 on direct  $R_s$  observations (light blue line) and their homogenized series (dark blue line)  
586 and sunshine duration (SunDu) derived  $R_s$  (light red line) and their homogenized series  
587 (dark red line). All of the lines were calculated based on observations at 41 sites. Details  
588 on how these 41 sites were selected are given in Section 3.1. The  $R_s$  variations are nearly  
589 the same as those shown in Figure 7, which were calculated based on all available  
590 observations.

591



592 1960 1965 1970 1975 1980 1985 1990 1995 2000 2005 2010 2015

593 Figure 7. Time series of annual anomalies of the surface incident solar radiation ( $R_s$ )  
594 based on direct observations (light blue line) and their homogenized series (dark blue  
595 line) and sunshine duration (SunDu) derived  $R_s$  (light red line) and their homogenized  
596 series (dark red line). All of the lines were calculated based on as many observations as  
597 possible. The light blue line and dark blue line were calculated from the  $R_s$  observations  
598 at 105 sites, while the light red line and dark red line were derived from the SunDu-  
599 derived  $R_s$  at 156 sites. The  $R_s$  variations are nearly the same as those shown in Figure  
600 6, which were calculated based on the 41 selected sites in Section 3.1. Large  
601 discrepancies were found in the homogenized data series (dark blue and dark red lines).  
602  
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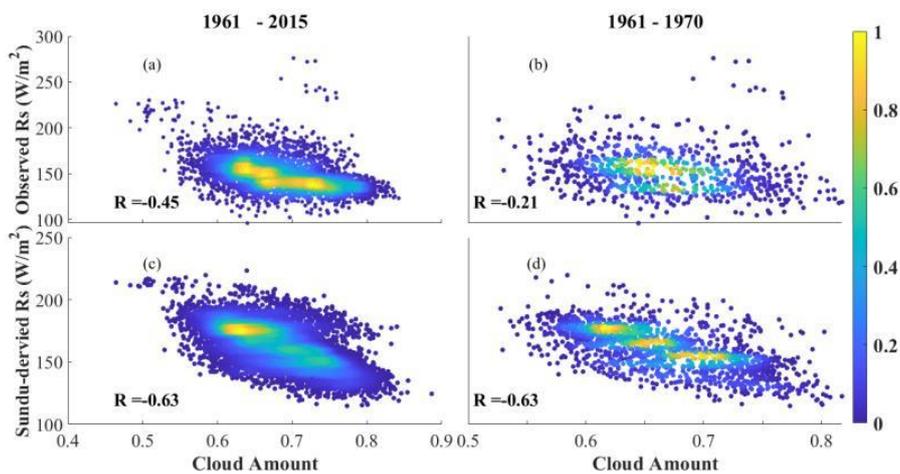
604

605 Figure 8. The cloud amount (CA) from CERES (blue line) agrees well with that derived  
606 from surface observations (red line) over Japan. At the annual time scale, the negative  
607 cloud radiative effect (-CRE, grey line) in CERES correlated well with the cloud  
608 amount.

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613 Figure 9. Scatter plot of homogenized monthly surface incident solar radiation ( $R_s$ )

614 (observed and SunDu-derived solar radiation) as a function of ground-based

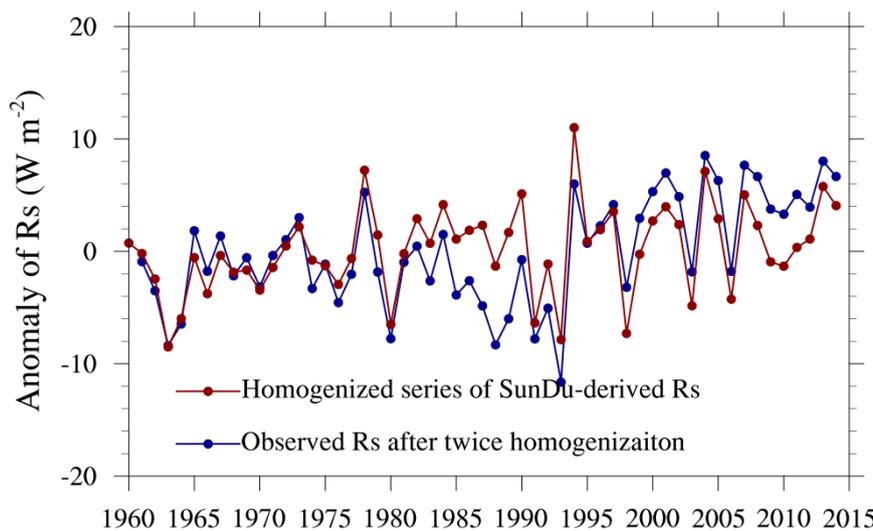
615 observations of cloud amount over Japan at all stations only when both cloud amount

616 data and observed  $R_s$  data are available. (a) and (c) for 1961-2015, (b) and (d) for

617 1961-1970. The smallest correlation coefficient in (b) indicates that the observed  $R_s$

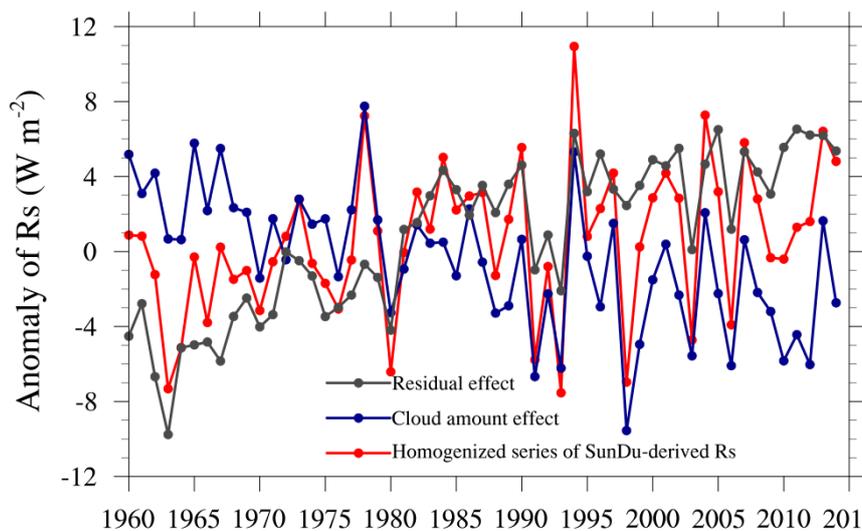
618 data are spurious for 1961-1970, and SunDu-derived  $R_s$  are more convincing.

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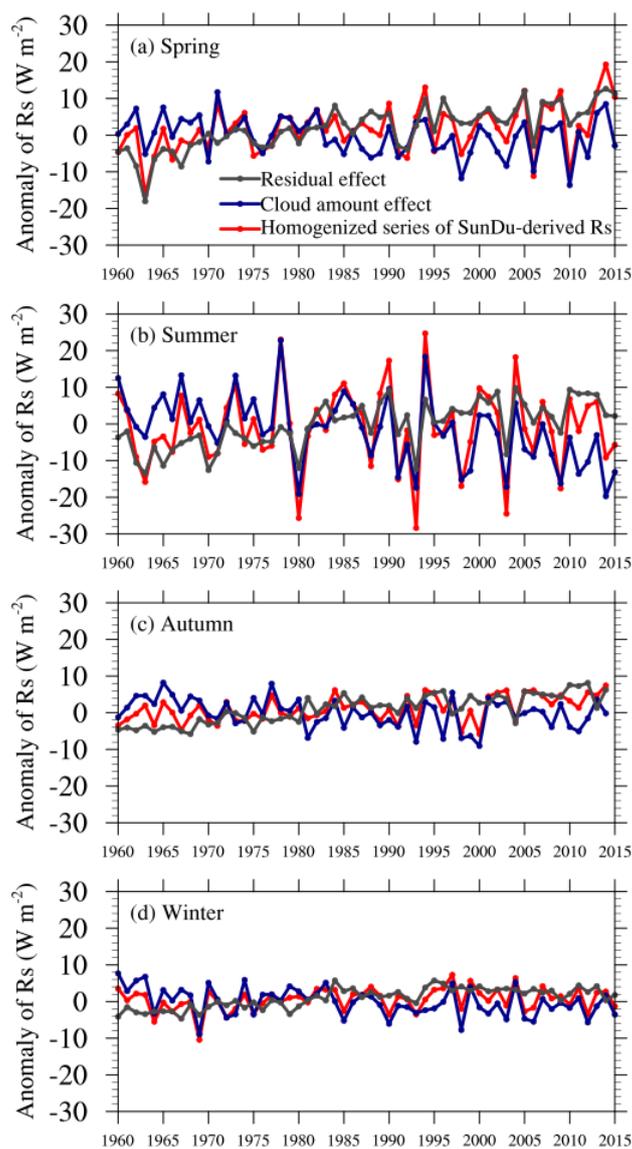
621 Figure 10. Time series of annual anomalies of the surface incident solar radiation ( $R_s$ )  
622 based on  $R_s$  observations after two homogenizations (dark blue line). The homogenized  
623 series of observed  $R_s$  from 1961 to 1970 shown in Figure 7 was tuned by RHtest method  
624 again using the homogenized series of SunDu-derived  $R_s$  (dark red line in Figure 7 and  
625 Figure 10) as a reference.



626

627 Figure 11. Area-averaged anomalies of homogenized SunDu-derived  $R_s$  (red line) over  
628 Japan. The cloud cover radiative effect (CCRE, blue line) was denoted as the change in  
629  $R_s$  produced by a change in cloud cover and calculated following Equation (4) by  
630 observed cloud amounts and cloud radiative effect (CRE) from the CERES satellite  
631 retrieval. The residual effect (grey line) was obtained by removing the cloud cover  
632 radiative effect (CCRE) from the homogenized SunDu-derived  $R_s$  anomalies.

633



634

635 Figure 12. Same as Figure 10 but for the four seasons. The decrease in Asian spring

636 dust may have triggered the brightening over Japan for 1961-2015, as the  $R_s$  in spring

637 increases most among the seasons.