



1 **High temporal resolution hydrometeorological data collected in the tropical**
2 **Cordillera Blanca, Peru (2004-2020)**

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26 **Abstract.** This article provides a comprehensive hydrometeorological dataset collected over the
27 past two decades throughout the Cordillera Blanca, Peru. The data recording sites, located in the
28 upper portion of the Rio Santa valley, also known as the Callejon de Huaylas, span an elevation
29 range of 3738 - 4750 m a.s.l. As many historical hydrological stations measuring daily discharge
30 across the region became defunct after their installation in the 1950s, there was a need for new
31 stations to be installed and an opportunity to increase the temporal resolution of the streamflow
32 observations. Through inter-institutional collaboration the hydrometeorological network
33 described in this paper was deployed with goals to evaluate how progressive glacier mass loss
34 was impacting stream hydrology, and to better understand the local manifestation of climate
35 change over diurnal to seasonal and interannual time scales. The four automatic weather stations
36 supply detailed meteorological observations, and are situated in a variety of mountain
37 landscapes, with one on a high-mountain pass, another next to a glacial lake, and two in glacially
38 carved valleys. Four additional temperature and relative humidity loggers complement the
39 weather stations within the Llanganuco valley by providing these data across an elevation
40 gradient. The six streamflow gauges are located in tributaries to the Rio Santa and collect high
41 temporal resolution runoff data. The datasets presented here are available freely from
42 <https://doi.org/10.4211/hs.059794371790407abd749576df8fd121> (Mateo et al., 2021).
43 Combined, the hydrological and meteorological data collected throughout the Cordillera Blanca
44 enable detailed research of atmospheric and hydrological processes in tropical high-mountain
45 terrain.

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57 **1 Introduction**

58 Glaciers and water resources in the Cordillera Blanca, Peru, have been under close
59 observation for nearly a century. In the 1930s, an Austrian geographer from Universitat
60 Innsbruck, Hans Kinzl, laid the groundwork for glaciological research in the region by surveying
61 and mapping the glaciers and identifying other natural features in the mountainous landscape
62 (Kaser and Osmaston, 2002). While the first systematic Peruvian effort to observe glacier tongue
63 variations in the Cordillera Blanca was initiated in 1944 by Broggi (Petersen et al., 1969),
64 glaciers as a source of security and water resources became the object of study in the subsequent
65 decades. The Unidad de Glaciología e Hidrología (later Unidad de Glaciología y Recursos
66 Hídricos (UGRH) of Electroperú S.A.) was initiated by the Corporación Peruana del Santa, a
67 government company for energy development, and took on glacier and lake monitoring after
68 many glacial lake outburst flooding (GLOF) events damaged communities downstream (Ames,
69 1998). GLOFs in 1941, 1945, and 1950 killed over 6000 people and destroyed a third of the
70 Ancash district capital of Huaraz (Carey, 2010; Carey et al., 2012). Other glacier hazards have
71 had detrimental impacts on the communities as well, including the 1970 avalanche, triggered by
72 a massive earthquake, which killed approximately 6000 people and along with the debris flow it
73 produced, covered the city of Yungay (Carey, 2010). Lliboutry, Morales and Schneider (1977)
74 investigated two glaciers in the mountain range in relation to the danger they presented for
75 flooding a downstream power plant. From the 1970s to the 1990s, a local Peruvian from Huaraz,
76 Alcides Ames, provided many important contributions to the present knowledge of glaciers in
77 the Cordillera Blanca while employed at UGRH, and after his retirement continued to dedicate
78 much of his time to studying the glaciers and sharing his knowledge with other researchers
79 (Ames et al., 1989; Hastenrath and Ames, 1995; Ames and Hastenrath, 1996; Ames, 1998; Kaser
80 and Osmaston, 2002). In the 1980s, Georg Kaser, an Austrian geographer, began studying the
81 Cordillera Blanca with a focus on the extent, causes, and possible consequences of the observed
82 glacier retreat (Kaser and Osmaston, 2002). For many decades, much of the research in the
83 region focused on developing better maps of the mountains and observing the marked glacier
84 retreat in the tropics. Although the meteorological and hydrological stations installed by UGRH
85 had been in place for nearly 40 years, it was not until the mid-1990s when studies began to
86 assess water resources in the Cordillera Blanca, but with minimal focus on flooding and GLOFs
87 (Kaser and Georges 1997; Mark and Seltzer, 2003).

88 The UGRH spent many years collecting an abundance of climatological, hydrological,
89 and glaciological data, which was useful for future researchers, hydroelectric and mining
90 companies, and other stakeholders throughout the Cordillera Blanca and its glaciers. Since the
91 mid-1990s, most of the daily-resolution discharge data from the Rio Santa watershed has been
92 cataloged and collected at a much reduced number of sites by private energy companies who
93 took over after Electroperu was privatized. The rapid turnover of the energy companies holding
94 the discharge observations makes it difficult to track down all available historical data. Also,
95 many of the hydrological stations have become defunct since they were installed in the 1950s,



96 creating a need for new stations to be installed for continuous monitoring of the region. A further
97 benefit of installing a new hydrological monitoring network in the 2000s was to significantly
98 increase the temporal resolution of the discharge observations.

99 Collaborating researchers from The Ohio State University, Bridgewater State University
100 and McGill University initiated a network of embedded environmental sensors in 2005. The
101 main goals of the instrument arrays were to evaluate how progressive glacier mass loss was
102 impacting stream hydrology, and to better understand the local manifestation of climate change
103 on the variability and controls of local weather phenomena over diurnal to seasonal and
104 interannual time scales. In order to create a sustainable network for continuous, long term
105 observation, this project has been maintained in close collaboration with Peruvian researchers
106 and government agencies, as well as with other international scientists to leverage resources in
107 maintaining instruments, in exchange for openly sharing data.

108 The instrumentation for the data collection we present in this work was installed and
109 maintained under a collaborative work agreement (“convenio”) formalized with the Peruvian
110 government agency overseeing the office in Huaraz (Peruvian Institute of Natural Resources
111 (INRENA) and the Autoridad Nacional del Agua (ANA)). This work agreement involved a
112 secondary collaboration with other international researchers who shared in installing and
113 calibrating the instrumentation. Specifically, Dr. Thomas Condom of the French Institut de
114 Recherche pour le Developpement (IRD) joined the agreement to install and maintain a series of
115 stream gauges logging water stage at 15 minute intervals. We also collaborated with the Austrian
116 research team of Dr. Irme Juen and Professor Georg Kaser from the Universitat Innsbruck,
117 Austria, who co-located a precipitation gauge with our weather station in Llanganuco. Further
118 details about the instrumentation are provided below.

119 Inter-institutional collaboration in this fashion has provided an effective partnership to
120 aid in maintaining the instrumentation over time, but also introduces challenges of logistical
121 coordination and data continuity. Visitation as non-resident international scientists to the field
122 sites has been feasible only one or two times annually. Thus, the Peruvians in our collaboration
123 have incorporated our instruments within their routine monitoring network. This has permitted
124 regular observations of stations and instruments to download data loggers, perform stream
125 discharge measurement to build a rating curve, and undertake a limited range of repair work.
126 However, limitations in local resources and manpower in Peru have often prevented recordings
127 of stage observation and discharge measurement to constrain rating curves. Likewise, having
128 multiple operators also increases some risk for data recovery errors. For instance, loggers that are
129 improperly relaunched after data downloads can jeopardize subsequent acquisitions. Having
130 more frequent site visits can allow for interventions, but also incurs increased risk of operator
131 error. Furthermore, Peruvian domestic political changes have disrupted the operations by
132 introducing different leadership, with altered operational priorities and resources, which directly
133 or indirectly interrupt the continuity of trained personnel responsible for data recovery and



134 preservation. Overcoming the many challenges of maintaining the instrumentation and
135 constructing rating curves has required regular cross-cultural communication, multiple and
136 annual visitation to the region.

137 Our collaborative observations have provided important new insights into how the
138 hydroclimate of the region is changing on different scales. Updated discharge and climate
139 observations in specific glacier catchments documented important shifts in seasonal supply of
140 glacier storage to the Yanamarey catchment (Bury et al., 2011) as well as suggesting regional
141 thresholds in glacier melt provision (Mark et al., 2010). The discharge constraints also provided
142 important validation for a novel hydrochemical basin characterization method to quantify
143 proportionate glacier melt and groundwater contributions to streamflow in tributaries of varying
144 glacierized coverage (Baraer et al., 2009, 2015). Regionally, the gauge network provided the key
145 constraint for a model of time-progressive hydrograph evaluation that verified significantly that
146 the main catchment had already passed “peak water” in the wake of strong glacier recession
147 underway for multiple decades (Baraer et al., 2012). Our embedded temperature and humidity
148 loggers distributed over elevation and linked to weather stations in the Llanganuco valley have
149 revealed novel diurnal to seasonal variations in lapse rates linked to catchment-specific valley
150 wind dynamics validated with downscaled climate models (Hellstrom et al., 2017). These studies
151 demonstrate the importance of collecting in situ hydrometeorological data and indicate the need
152 for continued data collection in the high Andes (Condom et al., 2020).

153 In this paper, we document available data from the Cordillera Blanca area collected over
154 the past 15 years. It is separated into (i) meteorological data recorded by permanently installed
155 automatic temperature and relative humidity loggers (Lascars), or automatic weather stations,
156 and (ii) hydrological data consisting of stage and discharge data from multiple sub-catchments.
157 The data are stored in .csv extension format on Consortium of Universities for the Advancement
158 of Hydrologic Science, Inc. (CUAHSI) Hydroshare:
159 <https://doi.org/10.4211/hs.059794371790407abd749576df8fd121> (Mateo et al., 2021). The data
160 collection consists of multiple time series of point observations from both meteorological and
161 hydrological measurement sites. These high-temporal resolution data provide detailed insight
162 into the complex hydro-meteorological system of the tropical Cordillera Blanca.

163 **2 Study Area - Cordillera Blanca**

164 The Rio Santa (Santa River) captures runoff from the western side of the glacierized
165 Cordillera Blanca and the eastern side of the non-glacierized Cordillera Negra, encompassing a
166 drainage area of 11636 km² at the outlet to the Pacific Ocean. While the headwaters of the Rio
167 Santa are found at Laguna Conococha at 4100 m above sea level (a.s.l.), the highest point in the
168 basin (and in Peru) is the summit of Huascarán at 6768 m a.s.l.. The average slope of the entire
169 drainage basin is 20.6° and its average elevation is 3374 m a.s.l. calculated from a 3 m resolution
170 DEM of the region. The Rio Santa flows over 300 km northwest from its origin at Laguna



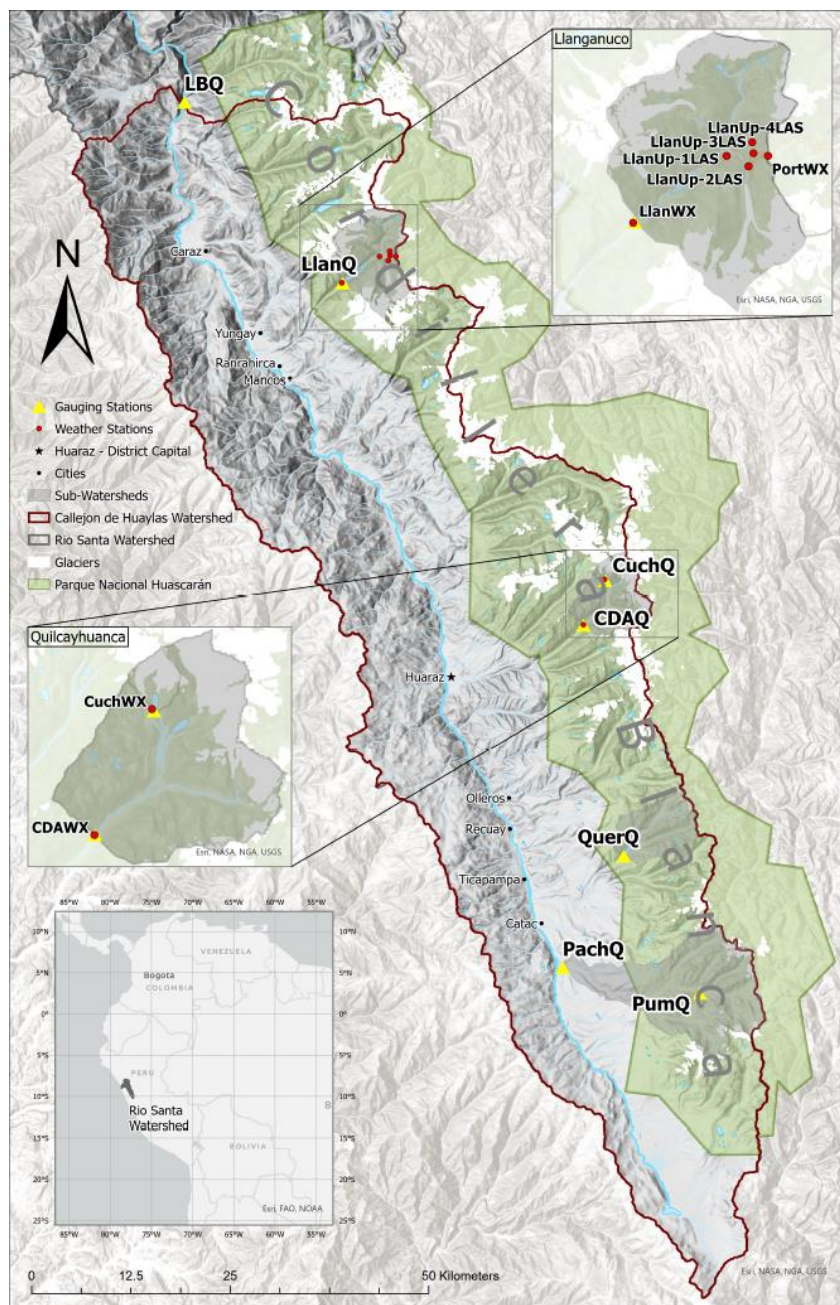
171 Conococha, an alpine lake at 4000 m a.s.l., to its outlet into the Pacific Ocean, near Chimbote.
172 The Callejon de Huaylas refers to the upper section of the Rio Santa, comprising approximately
173 4773 km² located above the Cañon del Pato 50 MW hydroelectric generation plant in Huallanca
174 (Figure 1). The Rio Santa basin is home to three other hydroelectric plants and provides water to
175 the expansive Chavimochic irrigation district near the coast (Mark and McKenzie, 2007). The
176 discharge on the Rio Santa has been carefully monitored since the Cañon del Pato dam began
177 operating in the 1950s, however, only one station near the dam remains active which is situated
178 slightly upstream at La Balsa (labeled as LBQ in Figure 1).

179 Figure 1.

180 Map of the Callejon de Huaylas showing the locations of the hydrological and meteorological
181 stations. (Base layers of the map originated from: © Esri, © NASA, © NGA, © USGS, © FAO,
182 © NOAA; all other layers were created and edited by authors of this article.)



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185 Rio Santa discharge experiences a strong seasonal contrast with the lowest discharge
186 occurring between July and September, while peak discharge, nearly 20 times greater,
187 occurs in March. Calculated from daily streamflow observations between 1954 and 2015, mean

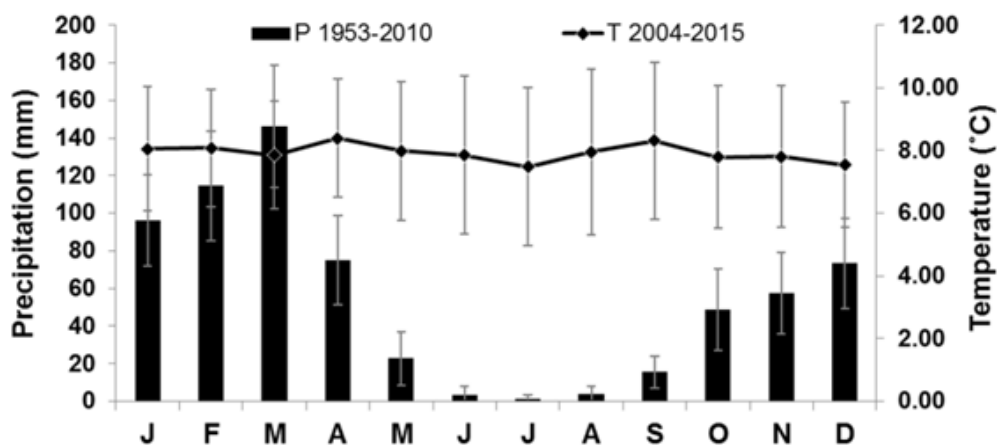


188 annual streamflow at La Balsa station on the Rio Santa, is $87 \text{ m}^3\text{s}^{-1}$, while the average annual
189 minimum discharge is $25 \text{ m}^3\text{s}^{-1}$, and the average annual maximum discharge is $445 \text{ m}^3\text{s}^{-1}$. The
190 Callejon de Huaylas is approximately 10% covered in ice (RGI Consortium, 2017), has an
191 average slope of 19° and an average elevation of 4055 m a.s.l.. Glacial melt in the Rio Santa at
192 La Balsa provides 10-20% of the total annual discharge and may exceed 40% in the dry season
193 (Mark et al., 2005; Condom et al., 2012).

194 The climate of the Callejon de Huaylas is semi-arid and displays distinct precipitation
195 seasonality, with the wet season between October and May being responsible for 80% of the
196 800-1200 mm/year of precipitation (Baraer et al., 2009), and the dry season lasting from June to
197 September (Figure 2). As a localized example, the Llanganuco valley displays an average total of
198 8 mm during the dry season, and an average of total of 258 mm during the wet season months of
199 December, January, and February, based on monthly totals from 1953 to 2010 (Hellstrom et al.,
200 2017). Variations in river discharge are largely driven by seasonality of precipitation.
201 Temperature remains nearly constant in the outer tropics, with the annual variation in
202 temperature being smaller than diurnal variation (Kaser et al., 1990).

203 Figure 2.

204 A climograph of the Cordillera Blanca displaying the strong contrast in precipitation between
205 wet and dry seasons, while maintaining a steady temperature throughout the entire year.



206

207 The geology in the Rio Santa basin is dictated by numerous tectonic and erosional
208 processes because the basin is situated along the active detachment fault of the Cordillera Blanca
209 (McNulty et al., 1998; Garver et al., 2005; Eddy et al., 2017). The highest peaks of the Cordillera
210 Blanca are composed largely of a granodiorite batholith intruded into a metamorphic unit that
211 includes hornfel, gneiss, and sulfide-rich lithologies, such as pyrite schist, phyllite, and pyrite-
212 bearing quartzite (Giovanni et al., 2010; Eddy et al., 2017). Granodiorite is found predominantly
213 in the northern portion of the range, while the southern portion of the range is made up of a



214 variety of metasediments, including quartzites and carbonates (Garver et al., 2005). The main
215 valley of the Rio Santa watershed is covered by recent sediment deposits, including alluvium,
216 landslide deposits, and glacial-fluvial fill. The impact of past and present glacier extent in the
217 topography is visible in geomorphic features throughout the range, including steep walled U-
218 shaped valleys, moraines, and proglacial lakes (Eddy et al., 2017).

219 **3 The Data**

220 In the following section, the data collected from the Cordillera Blanca are presented in
221 two collections. First, we provide a description of the meteorological stations and embedded
222 sensor network of Lascar data loggers, and then we present the time series data from the stations.
223 Second, we provide details of the setup and context of the discharge gauging station network,
224 and conclude by presenting the time series of discharge data and general statistics from them.

225 **3.1 Meteorological Data**

226 The Servicio Nacional de Meteorología e Hidrología del Perú (SENAMHI) has operated
227 a relatively dense national network of meteorological stations since 1964. Air temperature is
228 provided as daily T_{\max} and T_{\min} (additionally measured at 07:00, 13:00 and 19:00 local time), and
229 precipitation is measured once a day. In addition, the Universidad Nacional Santiago Antúnez de
230 Mayolo (UNASAM) has maintained a network of what was originally 16 meteorological stations
231 located at different elevations in the Cordillera Blanca (Ancash district) since 2012. Monthly
232 totals of precipitation data for the Rio Santa watershed have been collected by Electroperú South
233 America since 1953 although measurements were interrupted in 1994 with the privatization of
234 the respective institutions when most of the stations were abandoned and quality control was an
235 issue (Kaser et al., 2003). More recent installations of weather stations in the Cordillera Blanca
236 have been mostly associated with short-lived research projects typically lasting less than three
237 years (Hofer et al., 2010; Georges and Kaser, 2002). The hourly meteorological observations we
238 describe below were collected primarily from instrumentation we installed in two sub-
239 catchments of the Rio Santa drainage basin, Llanganuco, and Quilcayhuanca.

240 The Llanganuco valley is situated on the western side of the Cordillera Blanca across a
241 southwest ($\sim 240^\circ$) to northeast ($\sim 60^\circ$) axis with elevations ranging from 3400 m a.s.l. to 6746 m
242 a.s.l. Llanganuco is one of most glacierized valleys in the mountain range at 30% glacier
243 coverage, making it one of the most glacierized tropical valleys in the world. In the Llanganuco
244 catchment, our group has installed multiple weather stations dating back to 2007, although only
245 one remains active. This currently active station, labeled LlanWX, is located at 3835 m.a.s.l. near
246 the lower of the two largest valley lakes (Table 1).

247 Table 1.

248 This table provides general information about each meteorological station and lascar data logger
249 location within the embedded sensor network.



Station	Elevation (m a.s.l.)	Period of Operation	Lascar Error Adjustment	Slope Angle (°)	Slope Aspect (°)
Cuchillacocha - CuchWX	4642	2013-Present	--	0	NA
Casa de Agua - CDAWX	3924	2013-Present	--	0	NA
Llanganuco - LlanWX	3835	2007-Present	Yes	0	NA
Portachuelo - PortWX	4750	2006-2015	--	0	NA
LlanUp-1 LAS	3955	2006-2015; 2015-Present	No	14	301
LlanUp-2 LAS	4122	2006-2014; 2015-Present	No	33	259
LlanUp-3 LAS	4355	2006-2015; 2015-Present	No	31	236
LlanUp-4 LAS	4561	2006-2015; 2018-Present	Yes	33	225
Portachuelo LAS	4767	2007-2015	Yes	9	180

250

251 The embedded sensor network (ESN), as described by Hellström et al. (2010) and
 252 Hellström and Mark (2006), provides the in-situ meteorological data provided in the paper. In
 253 July 2004 the first automatic weather station (AWS), LlanWX, was installed near the lower lake
 254 in the Llanganuco valley to collect a continuous record of air temperature, wind speed, wind
 255 direction, relative humidity, solar irradiance, and soil temperature/moisture (Table 2). The AWS
 256 shown in Figure 3A is located in an open area on the valley floor and is surrounded by *Polylepis*
 257 trees. The site is protected in Huascarán National Park, and the location was previously used by
 258 the University of Innsbruck for precipitation measurements. The most significant wind
 259 obstructions are the steep bedrock walls of the valley toward the northwest and southeast which
 260 exceed 1000 meters above the valley floor. Northerly and southerly wind are occasionally
 261 recorded by LlanWX, and are likely caused by turbulence or lateral winds from the uneven
 262 heating of the valley walls. Winds flow parallel to the axis of the valley during approximately
 263 92% of the recorded time and are not greatly obstructed by surface vegetation. The sensors for
 264 LlanWX were originally sourced from the Onset Computer Corporation and logged with an
 265 Onset HOBO® data logger (<http://www.onsetcomp.com>) until 2014 these loggers were replaced
 266 with an Iridium® satellite Data Garrison logger (<http://www.upwardinnovations.com>). A 102-
 267 mm diameter radiation shield was used to reduce air temperature error caused by the sun except
 268 during the following years: 2007, 2011, 2012, 2013, and 2014; when a separate data logger was
 269 used as the primary source for air temperature and relative humidity observations. In 2013, our
 270 team upgraded the station by replacing the wind sensors with two new units from Onset and a
 271 new pyranometer from Apogee (<http://www.apogeeinstruments.com/pyranometer>) which
 272 exceeded the 1277 Wm⁻² maximum reading of the previous Onset sensors (Covert, 2016). The
 273 observations from LlanWX are largely continuous since 2005, with data gaps occurring
 274 occasionally between 2010 and 2014.

275 Table 2.

276 This table details the variables, sensors, and their accuracy which are collected at the
 277 meteorological stations throughout the region.



Variable	Sensor	Accuracy	Unit
Air temperature	Onset HOBO S-THB-M002 Temperature RH Smart Sensor	± 0.2 °C	°C
Precipitation	Onset HOBO S-RGB-M002 0.2 mm Rainfall tipping bucket Smart Sensor	± 0.2 mm	mm
Relative humidity	Onset HOBO Temperature RH Smart Sensor: S-THB-M002	± 2.5 %	%
Wind speed	Onset HOBO S-WSB-M003 Wind Speed Smart Sensor	± 1.1 m/s	m/s
Wind direction	Onset HOBO S-WDA-M003 Wind Direction Smart Sensor	± 5 °	°
Incoming solar radiation	Onset HOBO S-LIN-M003 Solar Radiation Smart Sensor	± 10 W/m ²	W/m ²
Atmospheric pressure	Onset HOBO S-BPB-CM50 Smart Barometric Pressure Sensor	± 3.0 mb	mb
Soil temperature	Onset HOBO S-TMB-M002 12-bit Temperature Smart Sensor	± 0.2 °C	°C

278

279 A second AWS, situated at Portachuelo (referred to as PortWX in Figure 1), was installed
 280 in July 2006 on a high pass at the top of the Llanganuco valley (4742 m a.s.l.). The station was
 281 situated on a steep, rocky ridge between the Llanganuco valley and the Vaqueria valley. There
 282 were only wind obstructions to the north due to a steep rock wall within 10 m of the station
 283 location. The AWS at Portachuelo had the same sensors as its counterpart, LlanWX, however
 284 was lacking an air temperature and relative humidity sensor with a radiation shield until it was
 285 installed in July 2015. Prior to this upgrade, these variables were recorded each hour by a data
 286 logger part of the ESN discussed in the following section. The Portachuelo station was stolen in
 287 2015 after recording data for nine years.

288 Figure 3.

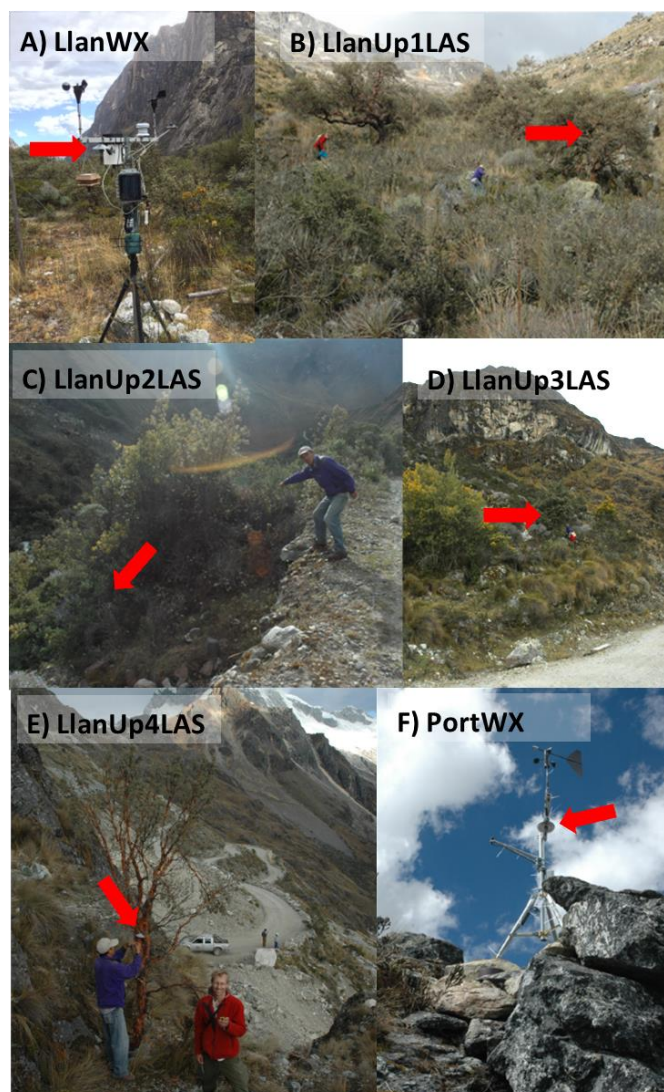


289 A) The typical setup for the HOBO automatic weather station, shown here is the LlanWX station
290 near the lower lake in Llanganuco. Measurements include: air temperature, humidity, wind speed
291 and direction, incoming solar radiation, rainfall, and soil temperature and soil moisture at -10cm.
292 Note also the Lascar in shield hanging from crossbar. B) Example of a Lascar data logger
293 location shaded from sunlight and hidden from view in a polylepis tree. Photographs taken by R.
294 Hellström (Covert, 2016).



295 A) 296 Smaller loggers measuring air temperature were installed in June 2005 to begin collecting
297 near surface temperatures to calculate lapse rates. These data loggers were a series of 8 nickel
298 sized iButton Thermochron® temperature loggers (Figure 3B). Hellström et al. (2010)
299 demonstrated the effectiveness of using the Thermochron loggers for purposes of observing near
300 surface lapse rates within the Llanganuco valley between the elevations 3470 and 4740 m a.s.l.
301 (Covert, 2016). The ESN of iButtons was replaced in July 2006 with a more robust network of
302 Lascar EI-USB2 data loggers (www.lascarelectronics.com) which were setup to measure air
303 temperature and relative humidity at one hour intervals. Figure 4 (from: Covert, 2016) provides
304 photos for a visual landscape context for each Lascar logger. While using the weather
305 observations provided, note that the LlanWX, LlanUp-4, and PortWX Lascar loggers are nearly
306 entirely exposed to sunlight which led to higher than expected air temperatures and greater error
307 from actual temperatures during the day (see section 4.1 for details). Each Lascar logger is
308 attached to a custom-designed and locally crafted radiation shield made of a thin tin cone and
309 two Styrofoam pieces in order to reduce error caused by direct sunlight, as shown in Figure 3B.
310 This Lascar network is still recording hourly data.

311 Figure 4.
312 Photos A-F show the locations of all Lascar dataloggers which gathered data provided in this
313 paper. Photographs taken by B. Mark and R. Hellström (Covert, 2016).



314

315 Quilcayhuanca contains two weather stations (Table 1) similar to the LlanWX station in
316 Llanganuco valley, the lower is located at 3924 m a.s.l. near an old river diversion station called
317 Casa de Agua (referred to as CDAWX in Figure 1), and the upper is located at 4642 m a.s.l.
318 slightly above an alpine lake, Cuchillacocha (referred to as CuchWX in Figure 1). The AWS at
319 Casa de Agua is still active since its installation in July 2013 and mostly continuous over this
320 duration. CDAWX station collects all of the same variables as LlanWX, which are described in
321 Table 2. The Cuchillacocha AWS was also installed in July 2013 and is still presently active, and
322 only has one data gap lasting longer than a month (from November 2016 – June 2017).

323 3.2 Hydrological Data

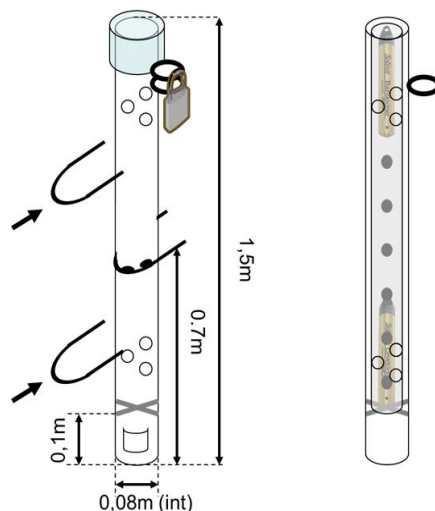


324 The La Balsa gauge (8.87° S, 77.82° W at 1880 m a.s.l.) (Duke Energy, Orazul Energy,
325 Inkia Energy) near the Cañon del Pato dam has a nearly complete record since 1954, acting as a
326 long-term reference point for discharge on the Rio Santa at the base of the Callejon de Huaylas
327 basin. Many other stream gauges and precipitation stations were put in place by the hydroelectric
328 companies, however, due to the lack of maintenance these stations became unusable. In 2008,
329 The Ohio State University, McGill University, IRD, and the Peruvian glaciology unit of ANA
330 commenced a joint project to reinstitute a stream gauging network throughout the Cordillera
331 Blanca (Baraer et al., 2012). Many of these redeployed stream gauges in sub-catchments of the
332 Rio Santa are still in working condition and are continuing to be monitored.

333 The stream gauges comprise custom-designed and locally crafted steel stilling wells
334 containing two Solinst leveloggers (Figure 5), with one measuring total pressure (water plus
335 atmospheric), and the other measuring barometric pressure. The Model 3001 Solinst levelogger
336 Edge has an accuracy of ± 0.05 kPa and a resolution of 0.002% Full Scale (FS) for pressure and
337 a temperature accuracy of $\pm 0.05^\circ$ C, and a temperature resolution of 0.003° C
338 (www.solinst.com). Water level is continuously monitored at these gauges at a temporal
339 resolution of 15-minutes by subtracting atmospheric pressure from the total pressure. Rating
340 curves at all stations were established and verified by conducting discharge measurements using
341 the velocity area method. Unavailable data is denoted by “NA” in the dataset. Linear
342 interpolation of missing data was only calculated for missing periods of less than 1 hour,
343 coinciding with the time the leveloggers were being downloaded in the field. Other than these
344 short periods of interpolated data, the datasets provided consist of raw data.

345 Figure 5.

346 Figure 5 illustrates the design of the stilling wells used to collect total pressure (water +
347 barometric) and barometric pressure.



348



349 The stream gauges that were installed in 2008 and 2009 and which are continuing to
350 collect discharge are Casa de Agua, Cuchillacocha, Pachacoto, Llanganuco, and Querococha
351 (Figure 1). The Pumapampa gauge went out of commission in 2016 but is also included in this
352 dataset (Table 3). These stream gauges have variable lengths of record due to periods of missing
353 data, but still provide valuable streamflow information in a region where stream records are
354 limited. As shown in the map (Figure 1), two pairs of these stream gauges are located at different
355 points along the same Rio Santa tributary in their respective sub-catchments: Pumapampa and
356 Pachacoto in the Pachocoto valley; and Cuchillacocha and Casa de Agua in the Quilcayhuanca
357 valley. All rating curves contain a minimum of eight point measurements throughout all times of
358 the year. The following two subsections provide a brief overview and summary for each gauging
359 station, organized by general location in the Cordillera Blanca, moving from the southern end of
360 the mountain range to the northern. Note that all glacier coverage is calculated from RGI 6.0
361 (RGI Consortium, 2017).

362 Table 3.

363 This table provides geographical information about each of the hydrological gauging stations.
364 Latitude and longitude are not provided for the security of the leveloggers. Note that all glacier
365 coverage is calculated from RGI 6.0 (RGI Consortium, 2017).

Station	Elevation (m a.s.l.)	Period of Operation	Contributing Area (km ²)	Percent basin covered in ice (%)
Cuchillacocha - CuchQ	4631	2008-Present	4.1	60.49
Casa de Agua - CDAQ	3948	2009-Present	66.9	24.78
Pumapampa - PumQ	4287	2008-2016	58.1	8.33
Pachacoto - PachQ	3738	2008-Present	202.3	4.95
Llanganuco - LlanQ	3850	2008-Present	86.9	30.32
Querococha - QuerQ	4005	2008-Present	62.9	1.53

367 3.2.1 The southern Cordillera Blanca

368 The Pumapampa catchment encompasses 58 km² with about 8% being covered by
369 glaciers (Table 3). Pumapampa gauge (referred to as PumQ in Figure 1) is located at 4287 m
370 a.s.l., below Nevado Pastoruri at the southern end of the Cordillera Blanca. This stream gauge is
371 situated in a channel through a low-lying meadow, where water overflowed its banks during
372 high-volume flows causing inaccuracies in high discharge measurements. Most notably, a series
373 of values in late-February indicated a peak discharge of 25 m³s⁻¹, well above extremes measured
374 from the rating curve. It was determined that during this period the leveloggers were overrun by
375 water and measurements became unreliable solely during this flooding event. Outside of this
376 event, discharge at Pumapampa is consistently within the measured values on the rating curve.
377 Mean annual streamflow (averaged from three years of data missing less than one-month during
378 the entire year) at this stream gauge was 2.3 m³s⁻¹ and a specific discharge of 1248 mm a⁻¹. Mean
379 streamflow during the dry season (May through September) in Pumapampa was 1 m³s⁻¹, while
380 the mean streamflow during the wet season (October through April) was 2.9 m³s⁻¹. Highest mean
381 monthly streamflow occurred in March 2014 with an average of 5.3 m³s⁻¹ (Figure 6).

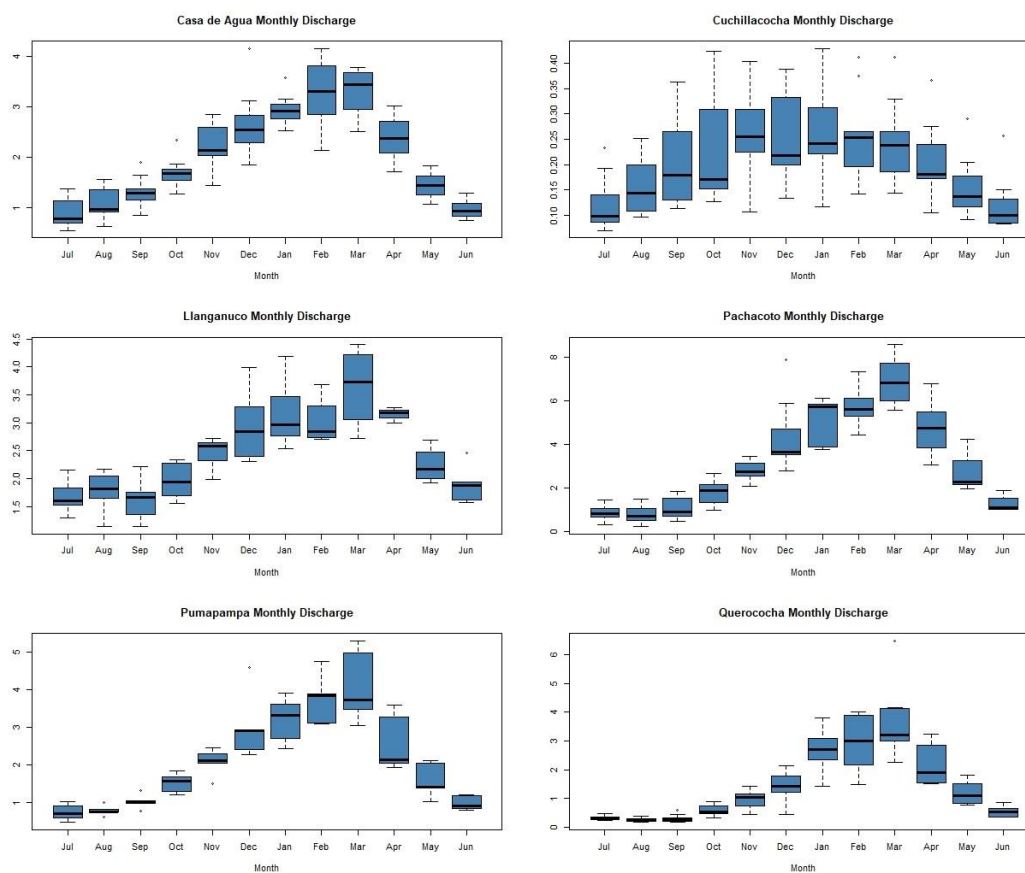


382 The Pachacoto gauge (referred to as PachQ in Figure 1) measures discharge 19 km
383 downstream from Pumapampa, at 3738 m a.s.l., just above its confluence with the Rio Santa.
384 Streamflow at Pachacoto is significantly greater than Pumapampa as it is the drainage point for a
385 larger area of 202 km², 5% of which is covered by glaciers. The mean annual streamflow at
386 Pachacoto was 3.2 m³s⁻¹, averaged from the seven near-complete years of data, with a calculated
387 specific discharge of 499 mm a⁻¹. Dry season runoff accounted for approximately 20% of the
388 annual streamflow by volume, while the majority of the remaining wet season runoff occurred
389 during February and March.

390 Streamflow at another sub-catchment in the southern Cordillera Blanca is recorded at the
391 Querococha gauge (referred to as QuerQ in Figure 1) at 4005 m a.s.l., located downstream of a
392 large lake and a drainage area of 63 km², including 1.5% of which is glacierized. This drainage
393 collects from two watersheds, one of which is entirely deglaciated, and one which has a rapidly
394 receding mass of ice. Using five years of complete data, the mean annual streamflow at this
395 gauging station was calculated to be 1.4 m³s⁻¹ and the specific discharge was 702 mm a⁻¹. The
396 streamflow from this gauge is somewhat regulated by the large lake residing 100 m upstream.
397 This basin is home to many previous studies from our research group, and many other
398 instruments throughout the basin are still consistently monitored by ANA.

399 Figure 6.

400 Figure 6 shows the monthly discharge at each of the gauging stations described in this study
401 calculated from complete months of data for each station. The strong seasonal pattern is clearly
402 visible at most of the stream gauges.



403

404 3.2.2 The central and northern Cordillera Blanca

405 Another catchment that contains two streamflow gauges is Quilcayhuanca, a valley
406 directly above the region's most populous city, Huaraz. The higher elevation stream gauge is
407 Cuchillacochoa (referred to as CuchQ in Figure 1) which is situated at 4631 m a.s.l. and measures
408 discharge below a high-alpine lake and two cirque glaciers encompassing a drainage area of 4
409 km² (with 61% of the basin covered in ice). This station displays a noticeably different discharge
410 pattern throughout the year than the lower gauging stations which collect greater runoff.
411 Cuchillacochoa discharge is not defined by a strong wet-dry season fluctuation, instead displaying
412 more variability each month and rising to peak values much earlier in the wet season than other
413 stream gauges. The average streamflow values at this location are also an order of magnitude
414 lower than other gauging stations due to the small drainage area it collects from (Figure 6).

415 The stream gauge located at a lower elevation in the Quilcayhuanca valley is Casa de
416 Agua (referred to as CDAQ in Figure 1) at 3948 m a.s.l. This gauging station is found near a
417 channel cut in the stream, in a large meadow, where water was previously rerouted for



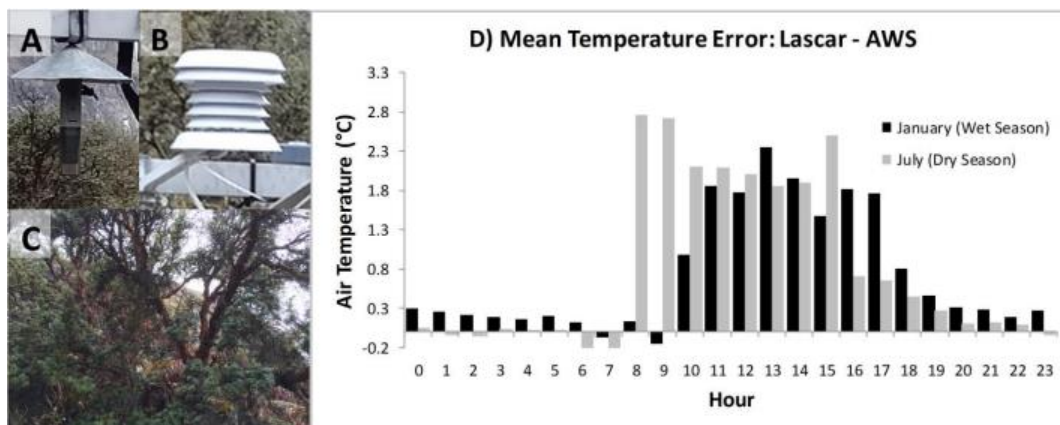
418 agricultural purposes. This gauge collects drainage from an area of 67 km² below the confluence
419 of two upper catchments, Cuchillacocha and Cayesh, well above the city of Huaraz (3100 m
420 a.s.l.). From this drainage point in the watershed, the basin is approximately 25% glacierized.
421 The mean annual streamflow, calculated from six complete years of data at Casa de Agua, was 2
422 m³s⁻¹, with a specific discharge of 943 mm a⁻¹.

423 The last streamflow gauge still recording data, Llanganuco (referred to as LlanQ in
424 Figure 1), is named after the valley it is located in and is positioned at 3850 m a.s.l., below two
425 valley lakes and a catchment area of 87 km², 30% of which is currently glacierized. The mean
426 annual streamflow, calculated from three complete years of data, was 2.3 m³s⁻¹, with a specific
427 discharge of 835 mm a⁻¹. A stream gauge has been in commission at Llanganuco on and off for
428 68 years, beginning as a gauge for UGRH and hydroelectric companies and later, after the
429 original station became defunct, a site for our new network of gauges.

430 **4 Data availability**

431 The datasets presented here are available freely from
432 <https://doi.org/10.4211/hs.059794371790407abd749576df8fd121> (Mateo et al., 2021). The
433 hydrological and meteorological data from the Cordillera Blanca have been used in short
434 segments in previous studies (Baraer et al., 2012; Baraer et al., 2015). Data availability varies
435 from station to station depending on location of loggers and instrumental errors which caused
436 periods of time to lapse without data being recorded. These datasets represent a majority of the
437 streamflow and meteorological data collected by our research group over the past two decades
438 collected at a high temporal resolution of 15-30 minutes. These include point discharge
439 measurements and short-time periods of 1-minute temporal resolution meteorological
440 measurements. The data provided here will be added to as future field seasons occur and the
441 effort to provide hydrometeorological data to the scientific community will continue at all of the
442 involved universities and institutions.

443 Figure 7.
444 Image A displays a Lascar data logger connected to its cone radiation shield. Two white
445 Styrofoam plates beneath the cone insulate the sensor from radiative heating. Image B shows the
446 radiation shields used on the AWS in comparison to the Lascar setup in image A. Image C shows
447 polylepis trees which are used to hide and shade Lascars in order to obtain accurate temperature
448 and relative humidity values. A comparison test between the Lascar setup and the AWS radiation
449 shield with no shading yielded the expected errors shown in graph D. Positive values indicate
450 that the Lascar reported a temperature higher than the AWS sensor. Photographs taken by R.
451 Hellström (Covert, 2016).
452



453

454 4.1 Radiation Error for Lascar Temperature Loggers

455 A year-long comparison between the Lascar setup (Figure 7) and LlanWX AWS
456 indicated an expected error which was largely dependent on the time of day. Variability in
457 temperature bias due to solar radiative heating is likely caused by changes in solar angle and
458 cloud cover patterns, particularly during the dry season. It is well documented that temperature
459 error rises on sunny days with calm wind and is reduced on windy and cloudy days regardless of
460 the season, but is far more prominent during the dry season in the tropical Andes (Georges and
461 Kaser, 2002). It is important to note that snow cover, which can create larger biases in
462 temperature by reflected solar radiation, is minimal throughout the year at the locations of the
463 Lascars in this study. Over an entire month, nightly error was within the 0.5 °C resolution of the
464 Lascar sensor. The occasional occurrence of greater nighttime error could be caused by the
465 sensor capturing longwave radiation emitted by the surface. Unlike the AWS radiation shield in
466 which the temperature sensor is fully enclosed, the Lascar is exposed from below meaning it is
467 able to capture longwave radiation at night. Further analysis will need to be done to verify this as
468 the cause of the nocturnal error, though these errors are considered minimal for the uses of these
469 datasets. The primary source of consistent instrumentation error for air temperature
470 measurements is the heating of radiation shields from those locations exposed to direct sunlight,
471 which required bias correction according to Figure 7. We bias-corrected for radiation error using
472 comparison with the LlanWX shielded temperature and Lascar logger hanging next to it under
473 the same exposure, including the LlanWX, LlanUp-4 LAS and PortWX station locations.
474 Because solar radiation is the most significant source of error, corrections were only applied to
475 daylight hours between 07:00 to 19:00 (Covert, 2016). During nighttime hours the error was
476 within the sensor's output resolution and no corrections were applied. In addition to corrections
477 for solar radiation heating, humidity and temperature records were assessed for accuracy and
478 removed if there were unrealistic deviations from the previous 24-hour trend of conditions within
479 the valley. Most of these deviations were due to an hour where sensors were downloaded and
480 held by hand or in a pocket, although data (less than 1%) were also removed if there was a



481 concern with the integrity of the sensor. Due to theft or replacement of radiation shields, the four
482 Lascar loggers were renamed to LanUp-1A, 2A, 3A, and 4A in June 2015 and this is reflected in
483 the database. The most common reason for larger gaps in data was either loss of sensors due to
484 theft or inability to access sensors because of logistical reasons, such as poor weather windows
485 or human error in downloading or redeploying sensors.

486 **5 Conclusions**

487 The Cordillera Blanca in the tropical Andes of Peru is a unique, high mountain region
488 with a wealth of high resolution meteorological and hydrological time series observations
489 collected over the past two decades. The region has been the focus of glacier monitoring efforts
490 for nearly a century. While daily time series of meteorological and hydrological observations
491 have been recorded off and on for nearly 70 years, there was no generalized sub-daily data
492 monitoring until the early 2000s. Maintaining this network of instruments recording high-
493 temporal resolution data involves traveling to remote areas of the Cordillera Blanca, protecting
494 the instruments from the harsh weather conditions at high elevations, and concealing instruments
495 to prevent theft. There is also difficulty in developing long-term strategies because most funding
496 agencies are focused on short-term based projects and do not fund continuous monitoring and
497 maintenance of data collection networks. With climate change as an ongoing global situation,
498 and its harsh impacts at local levels, long-term monitoring of hydrometeorological variables will
499 become of utmost importance to understand the changes that are occurring.

500 The datasets collected and described in this paper support investigations into climate
501 evolution, water resource availability, and hydrological changes across the Cordillera Blanca in
502 the past twenty-years. These datasets will also provide valuable, easy-to-access observations for
503 local water resource managers. This paper provides an overview of the variables which have
504 been measured and what is being made available as of 2021. Future measurements recorded in
505 the field will be made available as they are collected, to further build on the available hydro-
506 meteorological database in the Cordillera Blanca.

507 **Author Contribution**

508 EM prepared the manuscript with contributions from co-authors. EM, RH and MB curated the
509 data in preparation for this article. BGM initiated research investigation and acquired initial
510 funding for data collection network. All authors assisted in gathering data and maintaining the
511 network of data collecting stations.

512 **Competing Interests**

513 The authors declare that they have no conflict of interest.

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